

Sources and processes that control the submicron organic aerosol in an urban Mediterranean environment (Athens) using high temporal resolution chemical composition measurements.

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Abstract. Submicron aerosol chemical composition has been studied during a year-long period
(26/07/2016-31/07/2017) and two winter-time intensive campaigns (18/12/2013 – 21/02/2014 and
23/12/2015 – 17/02/2016), at a central site in Athens, Greece, using an Aerosol Chemical
20 Speciation Monitor (ACSM). Concurrent measurements include a Particle-Into-Liquid Sampler
(PILS-IC), a Scanning Mobility Particle Sizer (SMPS), an AE-33 Aethalometer and Ion
Chromatography analysis on 24 or 12-hour filter samples. The aim of the study was to characterize
the seasonal variability of the main submicron aerosol constituents and decipher the sources of
organic aerosol (OA). Organics were found to contribute almost half of the submicron mass, with
25 30-min resolution concentrations during wintertime reaching up to 200 $\mu\text{g m}^{-3}$. During winter (all
three campaigns combined), the primary sources contribute about 33% of the organic fraction,
comprising of biomass burning (10%), fossil fuel combustion (13%) and cooking (10%), while the
remaining 67% is attributed to secondary aerosol. The semi-volatile component of the oxidized
organic aerosol (SV-OOA; 22%) was found to be clearly linked to combustion sources and in
30 particular biomass burning, and even a part of the very oxidized, low-volatility component (LV-

OOA; 44%) could also be attributed to the oxidation of emissions from these primary combustion sources.

These results, based on the combined contribution of biomass burning organic aerosol (BBOA) and SV-OOA, indicate the importance of increased biomass burning in the urban environment of Athens as a result of the recession.. During summer, when concentrations of fine aerosols are considerably lower, more than 80% of the organic fraction is attributed to secondary aerosol (SV-OOA 31% and LV-OOA 53%). In contrast to winter, SV-OOA appears to result from a well-mixed type of aerosol, linked to fast photochemical processes and the oxidation of primary traffic and biogenic emissions. Finally, LV-OOA presents a more regional character in summer, owing to the oxidation, within a few days, of organic aerosol.

1. Introduction

Exposure to fine particulate matter is recognized as a leading cause of premature mortality in Europe. Even if the annual limit value is not exceeded at the majority of regulatory monitoring stations in European countries, health effects are expected to appear at lower levels as well, even below the WHO guideline values (EEA, 2017). Organic carbon (OC) is among the key PM components which record the strongest associations with short-term mortality (Ito et al., 2011; Klemm et al., 2011). Moreover, short-term exposure to OC has been linked to respiratory and cardiovascular hospital admissions (Levy et al., 2012; Zanobetti et al., 2009) and pediatric asthma emergency department visits (Strickland et al., 2010). In view of the health significance of fine aerosols, the characterization of their chemical properties and short term variability is critical, especially at the urban background level which is more relevant for the average population exposure. And while the majority of transformations related to particulate sulfate and nitrate have been well described, there is much progress to be made regarding the understanding of mechanisms that govern secondary organic aerosol (SOA) formation from precursors.

In this direction, the development of the Aerosol Mass Spectrometer (AMS) technology has been an important breakthrough, facilitating the study of aerosol chemical composition, in high temporal resolution. The ability to differentiate between primary and secondary components, based on specific markers, introduces an important advancement to organic aerosol (OA) source apportionment (Jimenez et al., 2009), which otherwise mainly relied on a statistical approach using elemental and organic carbon thermal-optical data (EC tracer method and variants; Turpin and

Huntzicker, 1995). Capitalizing on abundant spectroscopic data, PMF (Positive Matrix Factorization) source apportionment (SA) is used to discern between various primary sources like traffic and biomass burning, and to categorize secondary aerosols depending on their degree of oxidation. The ACSM (Aerosol Chemical Speciation Monitor) has been developed as an instrument relying on AMS technology while enabling long-term routine monitoring (Ng et al., 2011).

While many relevant studies have focused at regional and rural background areas, long-term ACSM results from large European urban centers are relatively scarce. Canonaco et al. (2013) have performed one year of measurements at an urban background site in the center of Zurich. Aurela et al. (2015) have deployed an ACSM at residential, traffic and highway sites within the Metropolitan Area of Helsinki, for a total of five months. Findings from 10 months of measurements at the North Kensington urban background site in London are reported by Reyes-Villegas et al. (2016). Focusing on Southern European cities, long-term results are provided by the intensive ACSM campaign of Minguillon et al. (2016), at an urban background site in Barcelona. Shorter - up to one month - studies using the AMS have been conducted in Barcelona (Mohr et al., 2012), Bologna (Gilardoni et al., 2016) and Marseille (El Haddad et al., 2013). In urban Athens, a one-month AMS campaign during winter 2013 has been carried out for chemical composition and OA sources (Florou et al., 2017).

The Greater Athens Area (GAA) appears as a challenging urban milieu for the study of aerosol dynamics, as it combines a large population (about 4 million) and intense primary emissions, with complex topography and meteorology, that lead to high levels of atmospheric pollutants and significantly deteriorate air quality (Kanakidou et al., 2011; Pateraki et al., 2014). However, the characteristics and related processes of secondary organic aerosols, in the long-term, have received up to this point limited attention (Grivas et al., 2012; Paraskevopoulou et al., 2014). Moreover, since 2013, due to the economic recession in Greece, primary and secondary precursor emissions have become altered and intensified, as the residents have switched from fossil fuel combustion to uncontrolled burning of wood and biomass for space heating (Saffari et al., 2013, Fourtziou et al., 2017, Gratsea et al., 2017). Existing measurements of aerosol chemical composition in Athens have been mainly performed using filter sampling (Theodosi et al., 2011, 2018; Paraskevopoulou et al., 2014) and have indicated the importance of fine organic aerosols. In this study we present, for the first time, long-term results on the sources of submicron organic aerosols in Athens from

high temporal resolution measurements during a year-long period, complemented by two intensive winter campaigns. For the collection of data, we deployed an Aerosol Chemical Speciation Monitor (ACSM) and in addition a Particle Into Liquid Sampler (PILS) coupled with ion chromatography, an AE-33 Aethalometer, while also conducting auxiliary aerosol (filter-based) and gas phase measurements. The main objectives are (i) to characterize the submicron aerosol and its variability using high temporal resolution, (ii) to quantify the sources of the organic aerosol and their seasonal variability (via PMF analysis) and (iii) to study the year-to-year changes of aerosol sources during winter time.

2. Experimental Methods

2.1 Sampling site and period

The measurements exploited in this study were conducted, at the urban background site of the National Observatory of Athens (NOA) at Thissio (37.97N, 23.72E), as representative of the mean population exposure over Athens metropolitan area (Fourtziou et al., 2017). The site stands at an elevation of 105 m above sea level, in a moderately populated area, where the influence of direct local emissions is limited.

The measurement period lasted for an entire year, from July 2016 to July 2017. Additionally, two intensive winter campaigns took place at the same site, the first from mid-December 2013 to mid-February 2014 and the second from 23 December 2015 to 17 February 2016. These intensive campaigns aimed at studying the year-to-year variability and impact of biomass burning on the air quality of the city of Athens during wintertime.

2.2 Instruments and Methods

Measurements were performed with an Aerosol Chemical Speciation Monitor (ACSM) by Aerodyne Research Inc. (Ng et al. 2011a), measuring the non-refractory PM₁ (NR-PM₁) chemical composition in near real-time (30-minute temporal resolution). The instrument was sampling through a BGI Inc. SCC 1.197 Sharp Cut Cyclone operated at 3 L min⁻¹, yielding a cut off diameter of approximately 2 μm. Practically, the ACSM operates following a similar principle as the Aerosol Mass Spectrometer (AMS) (Jayne et al., 2000) where ambient air is drawn through a critical orifice to a particle focusing aerodynamic lens; the resulting particle beam is flash-vaporized at 600°C, ionized via electron impact ionization and guided through a quadrupole mass

spectrometer. Ammonium nitrate and ammonium sulfate calibrations were performed prior to the
125 ACSM's deployment on the site for the period of 2016 – 2017 and the response factor (RF) for
nitrate along to the Relative Ionization Efficiencies (RIEs) for ammonium and sulfate were
determined. For the 2013 – 2014 and 2015 – 2016 intensive winter campaigns ammonium nitrate
calibration were performed and the RIE for sulfate was determined according to the fitting
130 approach proposed by Budisulistiorini et al. (2014). Values are presented in Table ST.1 of the
supplementary material. The detection limits for the ACSM provided by Ng et al. (2011a) are:
0.284 $\mu\text{g m}^{-3}$ for ammonium, 0.148 $\mu\text{g m}^{-3}$ for organics, 0.024 $\mu\text{g m}^{-3}$ for sulfate, 0.012 $\mu\text{g m}^{-3}$ for
nitrate, and 0.011 $\mu\text{g m}^{-3}$ for chloride. Mass concentrations are calculated using a chemical
composition dependent collection efficiency (Middlebrook et al., 2012) (Fig. SF.1).

Parallel measurements were performed for biomass burning identification, but also for quality
135 control purposes. In this context, a Metrohm ADI 2081 Particle Into Liquid Sampler (PILS; Orsini
et al., 2003) coupled with ion chromatography (Dionex ICS-1500) was used, which was sampling
ambient air from a different, but adjacent to the ACSM's, PM_{10} inlet. Two denuders were placed
inline, upstream of the instrument in order to remove gas phase species (e.g. NH_3 , HNO_3 , SO_2)
The ion chromatograph was set to measure cations such as ammonium and potassium at a time
140 resolution of 15 minutes. The resulting concentrations from the ACSM were tested against filter
measurements and the concentrations provided by the PILS. For the PILS, the detection limit was
calculated at 1 ppb for Na^+ , NH_4^+ and 2 ppb for K^+ . Non sea salt K^+ (nss- K^+) concentrations were
calculated using the Na^+ concentrations and the Na^+/K^+ ratio in seawater as a reference (Sciare et
al., 2005). Reported concentrations were blank corrected.

145 Furthermore, filter sampling was also conducted in parallel at the Thissio station. $\text{PM}_{2.5}$ aerosol
samples were collected on Quartz fiber filters (Flex Tissuquartz, 2500QAT-UP 47mm, PALL), on
a daily basis, while during the winter periods the sampling frequency was set to 12h. A
Dichotomous Partisol Sampler 2025 (Ruprecht & Patashnick) was used at a flow rate of 16.7 L
 min^{-1} . The samples were analyzed for organic and elemental carbon (OC, EC) with the Thermal-
150 Optical Transmission technique, using a Sunset Laboratories OC/EC Analyzer and applying the
EUSAAR-2 protocol (Cavalli et al., 2010). Filters were also analyzed for determination of the
main ionic species using ion chromatography as described in Paraskevopoulou et al. (2014).

Two different absorption photometers were monitoring Black Carbon (BC) concentrations. A
7-wavelength Magee Scientific AE-42 portable aethalometer was used for the 2013-14 and 2015-

155 16 winter campaigns, providing 5-min resolution measurements. For the year-long period a dual spot, 7-wavelength Magee Scientific AE-33 aethalometer (Drinovec et al., 2015) was used, operating at 1-min resolution and a 5 L min^{-1} flow rate. Standard gas analyzers for O_3 (Thermo Electron Co., model: 49i), CO , SO_2 and NO_x (HORIBA, 360 series) and a Scanning Mobility Particle Sizer for PM_{10} size distributions (SMPS 3034, TSI Inc.) measuring in the size range of 10.4
160 – 469.8 nm, were also operating at the sampling site. Wavelength dependent source apportionment of the BC load was performed by the AE-33 Aethalometer, based on the approach of Sandradewi et al. (2008) providing a fossil fuel (BC_{ff}) and a wood combustion (BC_{wb}) component. The default absorption Ångström exponents of 1 for fossil fuel combustion and 2 for pure wood burning, as incorporated in the AE-33 software, were used, very close to the respective values of 0.9 and 2,
165 used in a suburban site in Athens (Kalogridis et al., 2017). Meteorological parameters for the study were taken from the actinometric meteorological station of NOA, at Thissio (Kazadzis et al., 2018) (Fig. SF.2). All measurements were averaged to 1-hour intervals in order to synchronize the different data sets.

The bivariate wind speed-direction plotting methodology developed by Carslaw and Ropkins
170 (2012) in the Openair R-package, was used for the identification of source areas, as incorporated in the Zefir Igor Pro-based tool (Petit et al., 2017). Four-day back trajectories were also calculated using the HYbrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT_4) model (Draxler and Hess, 1998) developed by the Air Resources Laboratory (ARL/NOAA), and 1-degree GDAS (NCEP) meteorological data. Trajectories were computed every 3-h, for air masses arriving
175 at Athens at a height of 1000 m. The selected height is considered suitable to capture transport at a representative upper limit of the boundary layer in Athens (Markou and Kassomenos, 2010). Trajectory clustering was performed using the TrajStat plugin (Sirois and Bottenheim, 1995; Wang et al., 2009) of the MeteoInfo GIS software. The change of the total space variance for decreasing number of clusters was examined as a criterion for cluster number selection. The analysis was
180 performed separately for summer and winter, resulting in 5 clusters for each period.

2.3 Source apportionment of the submicron organic fraction using PMF analysis.

2.3.1 PMF strategy.

Positive Matrix Factorization (PMF; Paatero and Tapper, 1994) was performed on the organic
185 mass spectra obtained by the ACSM. The graphic interface SoFi (Source Finder) version 6.1,

developed at the Paul Scherrer Institute (PSI), Zurich (Canonaco et al., 2013) was used. SoFi implements the multilinear engine algorithm ME-2 (Paatero and Hopke, 2003), analyzing the acquired mass spectral timeseries matrix into a linear combination of factor profiles (FP) and time series sub-matrices. Detailed description of the method can be found in the above referenced studies.

For our datasets only $m/z \leq 125$ were used in order to avoid interferences from the naphthalene signal (m/z 127, 128 and 129). Weak signals, with signal-to-noise ratio (S/N) below 0.2 were downweighted by a factor of 10, and those with S/N between 0.2 and 1 were downweighted by a factor of 2 (Ulbrich et al., 2009), using the built in utilities of the SoFi toolkit.

The input organics and organics' error matrices are derived automatically from the ACSM data analysis software. Several model runs were performed, with and without applying constraints to the derived FPs, using the α value approach (Canonaco et al., 2013; 2015) and following the methodology proposed by Crippa et al. (2014). The α value ranges between 0 and 1 and is a measure of how much the resulting FPs are allowed to vary from the constraints. Initially, unconstrained PMF runs provided insight on the potential number and type of factors. For the following steps, reference factor profiles (RFPs) were introduced in order to constrain primary OA factors, (i) first for the Hydrocarbon – like organic aerosol (HOA), (ii) then for both HOA and BBOA and (iii) finally for HOA, BBOA plus cooking – like organic aerosol (COA). Potential FPs for secondary organic aerosols were left unconstrained. A thorough discussion on the choice and representativeness of the RFPs used can be found in section SI.4.1 of the supplementary material. Each factor was constrained using different α values within the limits suggested by Crippa et al. (2014). Next, the model's residuals, for each different model setup, were analyzed in search of structures that could indicate underestimation or overestimation of the number of separated factors. Stability of factors for different model seeds and correlations of the obtained FP spectra with FPs reported in similar environments and conditions were examined (Section SI.4.8). Finally, correlations of the time series of the selected optimal solutions to both gas phase and particulate independent measurements such as BC, BC_{ff}, BC_{wb}, CO, nss-K⁺, NO₃⁻, SO₄²⁻ and NH₄⁺ was examined to solidify the selection (Section SI.4.9).

The year-long data series was divided into a cold period, from November 2016 to March 2017 and warm period consisting of two sub-periods from August to September 2016 and from May to July 2017 which were treated separately. According to studies on the climatology of Southern

Greece, the transient period (spring and fall seasons) in Athens doesn't exceed 60 days on average (Argyriou et al., 2004), covering mainly the months of April and October - which were excluded from the seasonal analysis. The two wintertime campaigns of 2013-2014 and 2015-2016 were also treated separately.

The coefficient of determination r^2 for simple linear regression is used as a metric for all comparisons, e.g. both affinity of obtained FPs with literature spectra and correlation of the respective factor time series with independent measurements.

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2.3.2 Choosing the optimal configuration

Presentation of, and discussion on the optimal configuration chosen for the ME-2 model, as well as results from each step of the implemented strategy described above, followed by a sensitivity analysis on the α value influence on the obtained factors, can be found in section SI.4. In brief, for the cold period and the two winter – time intensive campaigns, constraining three factors, namely HOA, BBOA and COA, and leaving two unconstrained SOA factors, produces a solution that is characterized by minimal seed variability and model residual structures, while FPs, time-series, relative contribution and diurnal variability of the factors appear to be environmentally relevant, resembling solutions proposed earlier for the region (Kostenidou et al., 2015; Florou et al., 2017). Leaving factors unconstrained leads to an unstable model behavior such as diurnal residual structures for key variables (e.g. alkyl fragments like $m/z=55$ or 57) and large FP variability for different model seed runs. Furthermore, deconvolved spectra were missing expected variable contributions in profiles such as BBOA (very low $m/z=41$ and 43 relative contributions), while the COA factor was dominated by the CO_2^+ fragment at $m/z=44$. Configuring less or more than five factor solutions, resulted either in an even more pronounced residual diurnal cycle, pointing to poor factor separation or in splitting behavior and resulting factors which were environmentally irrelevant.

On the other hand, constraining two factors during the warm period, namely HOA and COA, and leaving two unconstrained SOA factors was found to be the solution exhibiting higher relevance while being robust and close to previous knowledge related to OA in the Greater Athens Area. A BBOA factor could not be identified during the warm periods, since contribution of the marker fragments for biomass burning $m/z=60$ and $m/z=73$ are almost absent in these periods

dataset. The COA factor is present in all the studied periods, validated following the approach of Mohr et al. (2012) (Fig. SF.9), and emerged in all the steps (unconstrained and constrained runs) of the implemented strategy (Figures SF.5 through SF.8 and related discussion in section SI.4).

3. Results and Discussion

3.1 Comparison of ACSM data with ancillary measurements

As a first quality control/quality assurance of the ACSM data, the ammonium concentrations are compared to the respective ones derived from the PILS, on an hourly basis for winter 2016-17. A good agreement is found ($r^2=0.80$, slope of 0.82). The sulfate and nitrate concentrations for the winter 2016-2017 period are compared to the respective ones from the ion chromatography analysis (PM_{2.5} filters), on a daily basis ($r^2=0.75$, slope of 0.81 and $r^2=0.78$, slope of 0.95, respectively). The organics concentrations are compared to the organic carbon concentrations of the PM_{2.5} filters. An excellent agreement is found ($r^2=0.93$, slope of 1.59) with the slope being close to values reported for urban areas (Petit et al., 2015) and OM:OC calculations from AMS measurements in polluted environments (Saarikoski et al., 2012). The results from the aforementioned comparisons are provided in the Supplementary material (SF.3).

During the intensive winter 2015-2016 campaign, the concentrations of the ACSM components are compared to those determined from the ion chromatography, based on concurrent filter samples collected at the same site, twice per day, (06:00 - 18:00 pm and t 18:00 - 06:00 local time). Results indicate an excellent agreement for sulfate ($r^2=0.88$, slope of 1.0), ammonium ($r^2=0.82$, slope of 1.06), and nitrate ($r^2=0.88$, slope of 1.12) (Figure SF.4). During the intensive winter 2013-2014 campaign, the ammonium concentrations from the ACSM showed significant correlation with the respective ones from the PILS ($r^2=0.80$, slope of 0.81).

Finally, the sum of the ACSM component concentrations plus BC, measured with the 7-wavelength aethalometer, was compared with the mass concentrations determined by the SMPS since February 2017 at Thissio. The density used to convert volume distributions and consequently volume concentrations of spherical particles to mass concentrations, was obtained by applying the methodology of Bougiatioti et al. (2014) assuming that the aerosol PM₁ population was dominated by ammonium sulfate and organics and calculating the respective mass fractions time series based on the ACSM measurements. A density of 1.77 g cm⁻³ was used for ammonium sulfate and 1.3 g cm⁻³ for organics (Florou et al., 2017). The results obtained using a chemical dependent collection

efficiency to determine the ACSM derived mass concentrations, are portrayed in Figure 1 and
280 indicate excellent correlation ($r^2=0.89$) a slope of 0.96 and an intercept of 0.60.

3.2 PM₁ average chemical composition and temporal variability

3.2.1 Chemical composition and characteristics

The time series of the main submicron aerosol components measured by the ACSM and the
285 black carbon concentrations are presented in the upper panel of Figure 2 (one complete year
period). The period average cumulative concentration of the ACSM components and BC was
 $12.4 \pm 12.5 \mu\text{g m}^{-3}$. The highest concentrations were measured during winter (average $16.1 \pm 19.5 \mu\text{g m}^{-3}$)
and the lowest during summer (average $10.3 \pm 5.6 \mu\text{g m}^{-3}$). On an annual basis, the most
abundant component was organic aerosol, followed by sulfate, contributing 44.5% and 27.8% to
290 the total submicron mass, respectively, while BC contribution was calculated at 15.1%, ammonium
at 7.9% and nitrate at 4.3%. In the middle and bottom panels of Figure 2 the respective time series
of the main submicron aerosol components during the two intensive 2-month winter campaigns
are presented. During winter 2013-14 the average mass concentration of the ACSM components
(plus BC concentrations) was $24.5 \pm 24.7 \mu\text{g m}^{-3}$, with organics and BC contributing 55.6 and 14.6%
295 to the total submicron mass, respectively, followed by sulfate (13.6%). During winter 2015-16 the
average concentration was $21.2 \pm 27.4 \mu\text{g m}^{-3}$, with organics and BC contributing 51.6 and 15.2%
to the total submicron mass, respectively, followed by sulfate (14.8%), nitrate (6.5%) and
ammonium (6.7%). It is clearly deduced that during the last winters, organics constitute half or
even more of the total PM₁ mass, sulfate around 20% and BC around 14%.

300 The other striking feature is that during wintertime, PM₁ concentration spikes can reach up to
 $220 \mu\text{g m}^{-3}$ hourly values, with organics taking up most of the mass. Maxima are recorded during
night-time and mostly during meteorological conditions favoring pollutants emission and
accumulation, such as low wind speed and low temperature (Fourtziou et al., 2017). There are on
average 8 such incidents occurring during each winter (10 in 2013-14, 7 in 2015-16 and 7 in 2016-
305 17), with organic levels being higher than $100 \mu\text{g m}^{-3}$. To our knowledge, such levels are the
highest reported for Europe during wintertime and highlight the strong impact of local emissions
and especially those related to heating/wood burning (see below), on the levels of organics and
consequently PM₁. Similar maxima to the ones observed in this study are also reported by Florou
et al. (2017, same site from 10 January until 9 February 2013), where organics concentration alone

310 reached up to $125 \mu\text{g m}^{-3}$ while maxima of $8 \mu\text{g m}^{-3}$ for BC and up to $5 \mu\text{g m}^{-3}$ for nitrate, were recorded. Similarly, wintertime pollution events with increased local character and elevated organics concentrations (up to around $100 \mu\text{g m}^{-3}$, average of $22.6 \mu\text{g m}^{-3}$) have been reported at a regional background site, just outside of Paris, during February 2012 (Petit et al., 2015).

315 **3.2.2 Seasonal variability**

The seasonal variability of the main measured species, along with the average PM_{10} concentration ($\mu\text{g m}^{-3}$), as calculated from the ACSM+BC measurements is shown in Figure 3 and the basic statistics are included in Table 1. Organics contribute 46% to the total submicron aerosol mass in summer, followed by sulfate (30.5%), BC (12.6%), ammonium (8.3%) and nitrate (2.6%), 320 while in winter, organics and sulfate contribute 48.1% and 23.2%, respectively, followed by BC (14.7%), ammonium (6.9%) and nitrate (6.3%).

The mass concentrations of organics, nitrate, chloride and BC exhibit a clear annual cycle, with minimum during summer and maximum in winter. This pattern seems to be due to a combination of three simultaneous processes. At first, the additional primary emissions from domestic heating 325 play an important role, as is evident by the largely elevated concentration levels of organics and BC, which during winter are also emitted by central heating systems and fireplaces. A second reason could be the decreased boundary layer depth during winter. According to Kassomenos et al., 1995 and Alexiou et al., 2018, daytime PBL depth shows a clear annual cycle, with maxima during the warm months (June to September) and exhibiting a two-fold decrease during 330 wintertime. Finally, the effect of temperature to the partitioning of the semi-volatile inorganics and organics can also contribute to the processes leading to the observed pattern. In support of the above, larger standard deviation is found in winter, demonstrating the frequency and magnitude of the observed pollution events due to the increased need for heating purposes (Fourtziou et al., 2017). Independently of the year, it can be seen that winter concentrations of organics, nitrate, 335 chloride and BC are very similar and more than twice the respective ones during the rest of the seasons (Table 1).

Organics concentrations are consistently high during all studied winters (from December to February), while the higher nitrate values, exhibiting similar trend with organics and BC can be attributed to the combination of lower temperatures during night-time along with the increased 340 combustion sources which lead to reduced acidity, resulting at the favorable partitioning of nitrate

in the aerosol phase (Park et al., 2005; Mariani and de Mello, 2007; Guo et al., 2016). Ammonium and sulfate exhibit the opposite seasonal cycle, with maximum values in summer and minimum during winter and spring. The higher summer sulfate levels are the result of enhanced photochemistry associated with more intense insolation, combined with less precipitation, favoring the regional transport of polluted air masses (Cusack et al., 2012). The seasonal variation of concentrations is in agreement with that observed in Athens, during prior long-term measurement campaigns based on analysis of daily filter samples (Theodosi et al. 2011, Paraskevopoulou et al., 2014; 2015).

3.2.3 Diurnal variability

When investigating the diurnal patterns of the measured species (Figure 4), it is observed that during wintertime, ammonium and sulfate do not exhibit significant variability, which is due to the regional character of ammonium sulfate. In order to quantify the extent of this variability we calculated the normalized diurnal pattern by dividing each hourly value with the respective species daily mean concentration. More specifically, sulfate varies by 13% around the mean value while ammonium varies by 40%. On the other hand, organics, BC and nitrate vary significantly during the day (183%, 79.8% and 110% respectively). These species clearly double their concentrations during night-time, caused by the additional primary emissions. Furthermore, BC also exhibits a second maximum during early morning hours, which should be attributed to the primary emissions during the morning traffic rush hours.

During summer, all concentrations are significantly lower, especially organics (note scale change) which exhibit a 5-fold decrease of their mean maximum concentration during night-time. Normalizing the diurnal cycles, as mentioned above, reveals a much less pronounced variability for organics (65%), implying a more regional character, while BC and nitrate exhibit the highest variability (67.7% and 77% respectively) in accordance to their local nature. The night-time maxima of BC vanish, while nitrate shows much lower concentrations, due to nitrate partitioning between gas and aerosol phase, favoring the vaporization of ammonium nitrate. BC still exhibits only one maximum during early morning hours owing to traffic emissions. Ammonium and sulfate diurnal profile follows expected photochemistry patterns, with peaking concentrations around 14:00 LT (UTC+02:00), consistent with secondary aerosol formation and increased vertical mixing with regional aerosol from aloft due to the evolution of the convective boundary layer

which exhibits a bell shaped diurnal structure ranging from a few hundred meters to above one kilometer, with maximum heights during early afternoon (Asimakopoulos et al., 2004; Tombrou et al., 2007). Finally, organics concentrations are somewhat higher during early night which could possibly be associated with biogenic/vegetation sources either local or regional that produce volatile compounds and condense on the particulate phase during night when temperatures are lower, as is further elaborated during the source apportionment results discussion in section 3.3. Furthermore, the organics variation also follows the late afternoon peak also observed for ammonium and sulfate. Condensation of the particulate phase could apply for nitrate as well, which also exhibits higher concentrations during night-time (almost double).

3.3 Source apportionment of organic aerosol

Warm period: In this period, the selected solution stems from a two factor constrained run (HOA using $\alpha = 0.05$ and COA using $\alpha = 0.1$) and consists of four factors: HOA, COA, SV-OOA and LV-OOA. As already mentioned, the two summer periods have been treated separately, but the derived spectra were almost identical (r^2 ranging from 0.98-0.99). The time series of the four identified sources during summer 2017 is shown in Figure 5 along with their diurnal variability and the respective average hourly contribution. The mass spectra of the selected solution are also provided in the supplementary material (Fig SF.12). No primary biomass burning aerosol could be identified, which is justified by the absence of fresh emissions over the city center during the warm period. In the summer periods HOA makes up 4.3% of the total organic fraction, while COA around 10% on average (7.3 and 11.3% for 2016 and 2017, respectively). In summer 2016 SV-OOA made up 32% and the rest 56% is LV-OOA. In summer 2017, SV-OOA contributes 34.6% to the total organic fraction while LV-OOA 49.7%. The dominance of secondary influence (SV-OOA & LV-OOA) is apparent, and accounts for the majority of the organic aerosol. This finding is in accordance with Kostenidou et al. (2015), who reported that 65% of the sampled aerosol during summer can be attributed to SOA (SV-OOA & LV-OOA), at a suburban site in Athens.

A comparison of the derived FPs with mass spectra in literature is shown in Figures SF.15 through SF.19 in the supplement. COA FP exhibits excellent correlation with spectra obtained in previous studies in the city (Florou et al., 2017; Kostenidou et al., 2015) as well as with spectra obtained in laboratory experiments investigating fresh OA emissions from meat charbroiling

(Kaltsonoudis et al., 2017). When calculating the O:C ratio in COA following the study of Canagaratna et al. (2015) we find a ratio of 0.19, which is comparable to the value of 0.24 obtained for COA during summer at a suburban site in Athens (Kostenidou et al., 2015).
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The HOA FP exhibits excellent correlation with literature spectra measured in cities located in the Mediterranean environment (Florou et al., 2017; Kostenidou et al., 2015; Gilardoni et al., 2016) as well as in other environmental and socioeconomical settings (Crippa et al., 2013; Lanz et al., 2009). According to Figure SF.18, where the affinity of SV-OOA with literature spectra is assessed, some assumptions could be made regarding the origin of the factor in this study. Similarity with Isoprene-Epoxydiol organic aerosol (IEPOX – OA), which is the oxidation product of isoprene, could denote a possible link of SV-OOA with biogenic aerosol. This association is further strengthened by considering the excellent correlation with SOA from biogenic precursors, such as a- and b-pinene reported by Bahreini et al. (2005) (r^2 of 0.86 and 0.89, respectively). These precursors are found to exhibit maxima during night-time (Harrison et al., 2001; Li et al., 2018; Hatch et al., 2011) coinciding with the diurnal behavior of SV-OOA in this study. On the other hand, comparison of the derived SV-OOA with SOA from diesel exhaust after 4 h of photochemical ageing (Sage et al. 2008) yields an r^2 of 0.89. Finally, SV-OOA exhibits the lowest correlations with the mass spectrum from aged organic aerosol emissions from meat charbroiling (Kaltsonoudis et al., 2017). The above mentioned comparisons with literature FPs, provide some indication that during summer, SV-OOA could be linked to SOA formation from the oxidation of volatile organic compounds (VOCs) from both biogenic and traffic sources and is not linked to the oxidation of primary COA. The low volatility component derived, exhibits excellent correlation to the very oxidized regional OOA found in the area (Bougiatioti et al., 2014) and good correlation with deconvolved OOA factors from previous studies in Athens (Florou et al., 2017; Kostenidou et al., 2015). When calculating the elemental ratios based on the study of Canagaratna et al. (2015), the O:C ratio for LV-OOA is 1.2, which is identical to the value of OOA obtained at Finokalia (Bougiatioti et al. 2014).
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In terms of comparison with independent measurements, HOA exhibits good correlation with nitrate ($r^2=0.62$) as well as with BC_{ff} ($r^2=0.63$), while COA, as expected, shows poor correlation with CO ($r^2=0.33$) and nitrate ($r^2=0.36$). SV-OOA is highly correlated with nitrate ($r^2=0.86$), implying common mechanisms in their variability, possibly linked with the partitioning between the gas and particulate phases. The poor correlation with CO ($r^2=0.4$) and BC ($r^2=0.35$) implies
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that SV-OOA may, to some extent, partially originate from a combustion source. LV-OOA shows
435 good correlation with sulfate ($r^2=0.62$) and ammonium ($r^2=0.63$), consistent with the regional
character of this factor. Results from the trajectory cluster analysis (Fig. SF.21) show that enhanced
LV-OOA levels are related to air masses originating from Eastern Europe and the Black Sea
region, which have both been identified as the main areas of influence for secondary aerosols that
are regionally processed and transported to Athens (Gerasopoulos et al., 2011; Grivas et al., 2018).
440 The regional character of LV-OOA is confirmed by high concentrations associated with increased
wind speeds (Fig. SF.20), especially those that originate from the Northern sector. These results
(presented in the Fig. SF.20 for the full dataset) are contrasted with HOA which displays a much
less diffuse spread, due to the intensity of local emissions (mainly traffic in the center of Athens).
The distant signal for LV-OOA to the SE direction could possibly be associated to processed
445 aerosol deriving from shipping activity (Petit et al., 2014) in the Aegean Sea.

Primary fossil fuel emissions (HOA) are very low during summer exhibiting a fivefold
decrease compared to the cold season, as in July and August most of the Athenians leave for their
summer vacations, thus reducing local traffic. Concentrations peak around 7:00 and after 19:00
LT that corresponds to the early morning and evening rush hours in downtown Athens. COA
450 exhibits a slight hump during lunch hours (13:00-15:00 LT) when concentrations rise to 65% of
the daily COA average after the morning minimum of around 50%, also seen in the relative
contribution of the factor, while a large night-time peak is present at around 22:00 LT. This late
peak, three times higher than the daily average value, is consistent with the late dinner hours and
operation of grill houses and restaurants in central Athens. SV-OOA exhibits 40% higher
455 concentrations during night-time compared to the SV-OOA daily average, which apart from
boundary layer dynamics may also be attributed to the condensation of semi-volatile compounds,
as also implied by the excellent correlation of the factor with nitrate. During daytime, following
the sharp decrease from the night time maxima, concentrations remain for some hours (10:00 to
14:00) close to 80% of the daily average before declining further in the afternoon. Finally, LV-
460 OOA exhibits a peak during mid-day, consistent with increased photochemical processes during
the peak of solar radiation intensity (Fig. SF.2) that lead to further organic aerosol oxidation.

In summary, during the warm period, the vast majority (more than 80%) of organic aerosol in
the area is linked to secondary organic aerosol formation. The semi-volatile product is of mixed
origin, linked to quick atmospheric processes, within a few hours, such as photochemistry of

465 primary sources, like biogenic emissions from vegetation, traffic emissions, or probably to a lesser
extent regional biomass burning. This last assumption could be supported by the fact that OOA
linked to aged BBOA has been reported in regional background sites in Greece (Bougiatioti et al.,
2014) and elsewhere (Minguillon et al., 2015), as well as by the fact that during the warm season,
470 air masses which mostly originate from the north, northeastern sector, carry pollutants from the
Balkans and around the Black Sea, areas heavily impacted by wildfires from July to September
(Sciare et al., 2008) (Fig. SF.21). On the contrary, the low-volatility product is the result of more
extensive oxidation of organic aerosol in the area, within a few days, and thus exhibits a more
regional character.

475 **Cold period:** In this period, the selected solution stems from a three factor constrained run
(HOA using $\alpha = 0.1$, COA using $\alpha = 0.2$, BBOA using $\alpha = 0.4$) and consists of five factors: BBOA,
HOA, COA, SV-OOA and LV-OOA. The solution for winter 2016-17 is presented (Fig. 6), while
the respective solutions for winter 2013-14 and 2015-16 are provided in the supplementary
material (Fig. SF.13). The time series of the five PMF factors for winter 2016-17 are shown in
480 Figure 6 along with their diurnal variability and the hourly contribution of each factor.

In terms of affinity with RFPs found in the literature, HOA for the cold season in this study is
found to exhibit excellent correlations with spectra obtained during the same season in earlier
studies in Athens as well as other Greek cities (e.g. Patras) (Florou et al., 2017) and also with HOA
factors obtained in different environments, a fact also observed for the warm season obtained
485 spectrum (Fig. SF.15). COA is excellently correlated with COA from Florou et al. (2017) in both
Athens and Patras as well as with COA measured by Kaltsonoudis et al. (2017) (Fig. SF.16). When
calculating the elemental ratios based on the study of Canagaratna et al. (2015) the O:C ratio for
COA is 0.18, which is in accordance with the value of 0.11 derived for COA at the same site by
Florou et al. (2017). BBOA exhibits high correlation with factors from Zurich, Paris and Finokalia
490 as summarized in Figure SF.17, while excellent correlation is found when compared to BBOA
found in Bologna, earlier studies in Athens and in Patras (Gilardoni et al., 2016; Florou et al.,
2017). Once more, the calculated O:C ratio for BBOA is 0.25, which is in accordance with the
value of 0.27 derived for BBOA at the same site by Florou et al. (2017). The SV-OOA spectrum
exhibits high correlation with the average SV-OOA from Ng et al. (2011b), as well as with the
495 IEPOX-OA from Budisulistiorini et al. (2013) ($r^2=0.80$ in both cases), as isoprene main oxidation

products such as methyl vinyl ketone and methacrolein are often used as biomass burning tracers (Santos et al., 2017). Similar correlation is also found with IEPOX-OA and SV-OOA during the winter 2015-16 campaign. The factor exhibits high correlation with SV-OOA from wintertime in Paris (Crippa et al., 2013) and SV-OOA from Hyytiälä (Äijälä et al., 2017) (Fig. SF.18). Finally, LV-OOA records an excellent correlation with the LV-OOA from Crippa et al. (2014), the average LV-OOA from Ng et al. (2011b), LV-OOA from Zurich during winter (Lanz et al., 2008) as well as with the oxidized OOA found in the region (Finokalia) (Bougiatioti et al., 2014) (Fig. SF.19).

The identification of BBOA is mainly based on the two fragments of m/z 60 and 73, considered as the “fingerprint” fragments of levoglucosan and biomass burning tracers. Indeed, BBOA exhibits excellent correlation with these two fragments ($r^2=0.94$ and 0.9 , respectively). $nss-K^+$ is also proposed as a very good tracer for biomass burning and as reported by Fourtziou et al. (2017), it shows a significant correlation with BC coming from wood burning (BC_{wb}), during wintertime in Athens. Consequently, the time series of $nss-K^+$ provided by the PILS-IC and m/z 60 are studied together. It appears that during both winters (2013-14 and 2016-17) for which $nss-K^+$ data is available, m/z 60 is in very good agreement with $nss-K^+$ ($r^2=0.85$) (Figure 7a). Furthermore, BBOA is highly correlated with BC_{wb} ($r^2=0.77$), and exhibits good correlation with $nss-K^+$ ($r^2=0.55$) and with CO ($r^2=0.51$). SV-OOA correlates excellently with both wood burning “fingerprint” fragments of m/z 60 and 73 ($r^2=0.99$ for both), highly with BC_{wb} ($r^2=0.90$) and CO ($r^2=0.73$) (Figure 7b) while exhibiting good correlation with $nss-K^+$ ($r^2=0.55$), demonstrating the direct link between SV-OOA and primary combustion sources (mainly biomass burning) (Table ST.2). It can be seen in Figure SF.21, that increased concentrations of both BBOA and SV-OOA are linked to air masses originating from northern and eastern Europe. During wintertime, these flow categories are associated with the prevalence of synoptic-scale northern winds and a decline in temperature in the area, leading to the appearance of PM episodes due to local combustion for residential heating (Paschalidou et al., 2015). The input of local sources confined in the Athens basin and in the vicinity of the sampling site is indicated by results of the wind analysis presented in Figure SF.20. Markedly enhanced levels are associated with weak or stagnant conditions. Results are contrasted with those of Grivas et al. (2018) at a moderately populated area in the eastern part of the basin. They found that local biomass burning emissions played a less important role than advections from the northern part of the area. In the present case, in the densely populated

center of Athens this effect is less apparent. The local character of wood burning aerosols in dense residential areas in Athens has also been indicated by Argyropoulos et al. (2017).

530 Comparison of the HOA time series with BC and CO yields a good correlation ($r^2=0.65$ and $r^2=0.65$ respectively). The factor correlates consistently better with BC_{ff} than with BC_{wb} (e.g. for the 2016 – 2017 r^2 is 0.60 versus 0.52, respectively). Correlation of COA with $nss-K^+$ and chloride ($0.3 < r^2 < 0.4$) could indicate a minor influence from emissions derived from biomass burning in meat-cooking (Akagi et al. 2011; Kaltsonoudis et al., 2017). Finally, LV-OOA showed a good correlation with ammonium ($r^2=0.58$), nitrate ($r^2=0.61$), $nss-K^+$ ($r^2=0.4$) and m/z 73 ($r^2=0.51$), demonstrating that part of the very oxidized OA during wintertime may also originate from
535 combustion sources as well.

Therefore, during the cold period, the organic aerosol in the area linked to secondary organic aerosol formation contributes around 65% to the total organic fraction. In contrast to summer, the semi-volatile products seem to be linked to the fast oxidation of primary combustion emissions (e.g. BBOA), which is also reflected on its diurnal variability (Fig. 6) and also on the strong
540 correlations with external tracers of primary combustion (see Table ST.2). Affinity to biomass burning tracers points out that the largest part of SV-OOA originates from the fast oxidation of BBOA. The low-volatility product is, in this case, likely of more local than long-range transport nature, as also highlighted by the almost two-fold higher values during night-time.

The diurnal cycles of the five factors are shown in Figure 6. HOA, originating from fossil fuel
545 combustion, exhibits maximum values during night, associated with combustion from central heating, and presents a secondary peak at 09:00 coinciding with the early morning traffic rush hour. The association of the factor to local primary emissions is also corroborated by the wind analysis plots (Figure SF.20). The dependence of HOA on wind speed and direction is similar between cold and warm seasons. The concentration vs. wind speed distribution, displays a wind
550 dilution effect and is characteristic for traffic-related fine particles in Athens (Chaloulakou et al., 2003; Kassomenos et al., 2012).

COA has similar winter and summer diurnal profiles with a moderate hump, with concentrations rising from 30% ($0.3 \mu g m^{-3}$) to 60% ($0.6 \mu g m^{-3}$) of the daily average ($0.98 \mu g m^{-3}$) during lunch hours (12:00-15:00 LT) and a large night-time peak (approx. 22:00 LT), partly
555 controlled by the decrease of the planetary boundary layer, but also owing to the expected increase in the activity of numerous restaurants in the area. A similar diurnal cycle for COA has been

reported by Florou et al. (2017). BBOA is characterized by a pronounced diurnal cycle with peaking values during night-time, associated with the production of this component in the evening by combustion for heating purposes. SV-OOA exhibits the largest diurnal amplitude, with night-time values being almost 6-fold higher compared to daytime. A plateau, with concentrations of SV-OOA being around 50% of the daily average value, following the sharp decline after midnight, is observed during the morning traffic rush hour, before another decline occurs until the daily minimum is reached at 14:00, demonstrating the possibility of the factor's provenance from the oxidation of freshly-emitted primary combustion organic aerosol. Finally, LV-OOA also exhibits 2-fold higher values during night-time compared to daytime. It has a similar to the SV-OOA factor behavior, with a secondary peak at 10:00, exhibiting a 1 – hour lag after the morning traffic rush hour, showing once more that part of the low volatility OA may also originate from the fast oxidation of primary combustion emissions, as also implied by its correlation with combustion tracers.

Table 2 sums up the contribution of each one of the 5 identified factors during the three studied winters. Overall, during wintertime BBOA constitutes around 10% of the total organic fraction. Based on the diurnal variability of this factor, its contribution is more pronounced during night-time, when concentrations are 4-fold or higher than the daytime ones, matching emissions from fossil fuel combustion represented by the HOA factor incorporating both traffic and heating oil combustion. Even though an exact mechanism has yet to be established, our assumption that the larger part of the SV-OOA comes from the rapid oxidation of freshly emitted BBOA through processes which involve nitrate radicals and/or heterogeneous reactions, appears justified via the excellent correlations with biomass burning tracers as well as by considering similar assessments found in other studies (Lathem et al., 2013; Cubison et al., 20011; Bougiatioti et al., 2014). In this manner the overall contribution of biomass burning becomes even more significant. Given that SV-OOA contributes around 30% to the organic mass, it is evident that during wintertime, biomass burning may contribute almost half of the total organic aerosol, with this contribution maximizing during night-time. More specifically, for BBOA the lowest contribution during daytime is 5.5% reaching a maximum of 27.5% during night (Figure 6). The same applies to SV-OOA with daytime minimum contribution of 13.8% and night-time maximum of 34.9%. What is also very important is the fact that even though the winter and summer mass spectra of SV-OOA have some similarities ($r^2=0.83$), there are also differences, especially in the origin of this component, as during winter

the majority is linked to the oxidation of primary combustion sources, while during summer the absence of a significant correlation with BC or nss-K⁺ implies the presence of different sources, both anthropogenic (but not biomass burning) and possibly biogenic.

4. Summary and conclusions

High temporal-resolution measurements were conducted for an entire year (plus two, two-month duration, intensive measurement campaigns during wintertime) at an urban background site in Athens, using an ACSM, a PILS-IC system and an aethalometer, in addition to routine pollution measurements. During the 16 – month measurement period, several pollution events with PM₁ concentrations reaching as high as 220 µg m⁻³ were recorded, all encountered during wintertime nights. In these cases, organics contributed the largest fraction to the submicron particulate mass, with overall contribution during wintertime reaching 50%, followed by sulfate (~20%) and BC (~14%). Within a typical winter day, organics, BC and nitrate double their concentrations during night-time. The increase of the first two can be attributed to emissions linked with domestic heating while nitrate exhibits higher concentrations due to the combined effect of decreased temperature and aerosol acidity, favoring partitioning in the aerosol phase. During summer, organics, BC and nitrate concentrations are significantly lower while sulfate and ammonium levels are increased. Organics are once more the main aerosol constituent contributing by 46%, followed by sulfate (30.5%), ammonium (8.3%) and BC (8%). Within a typical summer day, ammonium and sulfate concentrations peak at about 14:00 LT (UTC+2), consistent with secondary aerosol formation.

Organics, nitrate, chloride and BC exhibited a clear seasonal cycle with maximum during winter and minimum during summer. Sulfate and ammonium exhibited the opposite cycle, as a result of enhanced photochemistry, limited precipitation and higher regional transport.

Based on the source apportionment of the organic aerosol, four factors were identified during summer, namely HOA, COA, SV-OOA and LV-OOA, and five factors during winter, the same as in summer with the addition of primary biomass burning emissions (BBOA). During summer, HOA makes up 4.3% of the total organic fraction, COA around 10%, and the rest is linked to secondary organics (SV-OOA and LV-OOA). HOA has peak values during the morning traffic rush hour, and COA mainly during night-time. SV-OOA exhibits two-fold higher concentrations during night-time while LV-OOA exhibits a peak during mid-day, consistent with photochemical processes. The semi-volatile product is clearly of mixed origin, linked to quick

atmospheric processing within a few hours, of VOCs emitted from primary sources like vegetation,
620 traffic and to some limited extent to processed regional biomass burning. The low-volatility
product, on the other hand, is the result of more excessive oxidation, in the order of several days,
having thus a more regional character.

Combining the results from the three different winter campaigns, HOA accounts for almost
13% of the organic fraction, COA around 10%, BBOA 10%, SV-OOA 22% and LV-OOA 45%.
625 All constituents exhibit significantly higher concentrations during night-time, with HOA being
also linked to primary emissions by heating oil combustion from central heating units and
presenting a secondary peak during the morning traffic rush hours. COA has a similar diurnal
profile to the one observed during summer. BBOA is also characterized by a pronounced diurnal
cycle with peaking values during night from combustion for heating. SV-OOA has almost 6-fold
630 higher concentrations during night, consistent with its link to the oxidation of primary combustion
sources, while even LV-OOA exhibits almost 2-fold higher concentrations during night. In
contrast to summer, the semi-volatile product during winter has a very clear origin, linked to the
fast oxidation of primary combustion sources (HOA and BBOA) with BBOA being the major
source, due to the affinity of SV-OOA with biomass burning tracers. Part of the LV-OOA, as well,
635 could originate from the extensive oxidation of the local primary combustion sources, showing
that LV-OOA during winter is of more local than regional character.

Concluding, it is clear that organic aerosol constitutes a large fraction of submicron aerosol
throughout the year, in the urban environment of Athens. During wintertime, a large part of this
OA, as high as 50%, originates from combustion sources for heating purposes, such as biomass
640 burning and diesel oil fueled central heating, causing significant air quality deterioration. Night-
time contribution of BBOA is 7-fold higher than the one during day, while the respective
contribution of SV-OOA is increased by a factor of 2.6. Given that during wintertime, fine PM
concentrations reach up to $220 \mu\text{g m}^{-3}$, the significance of these sources contribution to air quality
degradation becomes even more striking, demonstrating the necessity for strategic, long-term
645 mitigation actions.

650 **Data availability**

All data related to the publication are available upon request from Nikolaos Mihalopoulos (nmihalo@noa.gr).

Competing interests

655 The authors declare that they have no conflict of interest.

Author contribution

NM, EG and AB conceived and facilitated the study. IS, AB, DP, EL, PZ and MT contributed to the measurements. IS, AB, GG and NM performed the analysis and wrote the paper. All authors
660 commented on the manuscript.

Acknowledgments

I. Stavroulas and N. Mihalopoulos acknowledge support by the State Scholarship Foundation
665 (“IKY Fellowships of Excellence for Postgraduate Studies in Greece -Siemens Programme, 2016-2017”), in the framework of the Hellenic Republic-Siemens Settlement Agreement. The authors would also like to acknowledge support from Francesco Canonaco and Andre Prévôt from PSI, who developed SoFi and provided valuable input related to Positive Matrix Factorization. This study contributes to ChArMEx work package 1 on emissions and sources.

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Table and Figure Captions

1020 **Table 1:** Seasonal average concentrations \pm standard deviation (range) and total mass of the main submicron aerosol components for the one-year study period and the two winter campaigns.

Table 2: Contribution of the five organic aerosol components to the total organic fraction during the three individual winter campaigns.

1025 **Figure 1:** Correlation between ACSM+BC vs. SMPS-derived 1 – hour averaged mass concentrations for the 2016-17 measurement period.

1030 **Figure 2:** Time series of the main submicron aerosol components. On the top panel the one-year period starting on 26 July 2016 and ending on 31 July 2017, on the middle panel the 2013-2014 winter campaign (18 December-21 February), and on the bottom panel, the 2015-2016 winter campaign (23 December-17 February).

1035 **Figure 3:** Monthly average concentrations of the main aerosol constituents. Organics are shown on the top panel for the one year 2016-2017 period as well as the 2013-2014 and 2015-2016 winter periods, while sulfate and ammonium on the middle panel, and BC, nitrate and chloride on the bottom panel shown for the one year 2016-2017 period. Standard deviation is also depicted (error bars; only the positive part is shown for plot's clarity issues).

1040 **Figure 4:** Average daily cycle of the main submicron aerosol constituents for the cold period 2016-17 on the top panel and the warm period of 2017 on the bottom panel. The size of the markers indicates the normalized values relative to each species' daily mean value.

1045 **Figure 5:** Time series of the contribution of the different factors identified by PMF between 1 May – 31 July 2017 (top) along with their average diurnal cycles (bottom left) and the respective hourly average contributions (bottom right).

Figure 6: Time series of the contribution of the different factors identified by PMF between 21 Nov. 2016 – 1 March 2017 (top) along with their average diurnal cycle (bottom left) and respective hourly contribution (bottom right).

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Figure 7: (a) Correlation of m/z 60 with nss-K⁺ for 2013-14 (red) and 2016-17 (blue), and (b) Correlation of SV-OOA with CO (grey) and BC (blue) for 2016-17.

1055 **Table 1**

	Mar – Apr – May 2017	Jul-Aug 2016 & Jun- Jul 2017	Sep – Oct – Nov 2016	Dec – Jan – Feb 2016- 2017	Dec – Jan – Feb 2013- 2014	Dec – Jan – Feb 2015- 2016
Organics	3.3 ± 3.0 (0.3-31.3)	5.4 ± 3.4 (0.3-41.9)	6.1 ± 7.5 (0.1-98.2)	9.0 ± 13.4 (0.2-153.9)	18 ± 24.4 (0.4-212.2)	12.4 ± 19.9 (0.7-1150.5)
Ammonium	0.6 ± 0.5 (0.4-3.1)	1.0 ± 0.6 (0.2-4.1)	1.0 ± 0.7 (0.4-5.7)	0.9 ± 0.7 (0.2-5.7)	1.8 ± 1.2 (0.2-9.1)	1.1 ± 1 (0.3-6.7)
Sulfate	2.1 ± 1.5 (0.2-10.1)	3.6 ± 2.1 (0.3-14.9)	3.5 ± 2.3 (0.1-17.1)	2.5 ± 1.5 (0.1-11.7)	2.6 ± 1.4 (0.4-13.9)	2.2 ± 1.7 (0.4-10.3)
Nitrate	0.4 ± 0.5 (0.05-5.4)	0.3 ± 0.2 (0.01-1.5)	0.5 ± 0.7 (0.1-6.9)	1.2 ± 1.5 (0.05-12.1)	2.6 ± 2.4 (0.09-18.3)	1.5 ± 1.4 (0.07-16)
Chloride	0.02 ± 0.05 (0-0.8)	0.02 ± 0.02 (0.04-0.2)	0.04 ± 0.09 (0.07-2.0)	0.15 ± 0.3 (0-3.5)	0.16 ± 0.24 (0.09-8.1)	0.12 ± 0.24 (0-2.6)
BC	1.5 ± 1.4 (0.1-14.6)	1.2 ± 0.8 (0.2-10.5)	1.7 ± 1.6 (0.1-12.4)	2.4 ± 3.4 (0.1-29.6)	2.7 ± 3.2 (0.2-26.8)	3.4 ± 4.6 (0.2-32.3)
PM1	8.9 ± 6.1 (0.6-42.4)	10.3 ± 5.6 (0.5-52.2)	13 ± 11.1 (0.9-115.5)	16.1 ± 19.5 (0.8-185.8)	24.5 ± 24.7(1.4- 227.2)	21.2 ± 27.4 (1.7-215.3)

1060 **Table 2**

	Winter 2013-14 18/12/13 – 21/02/14	Winter 2015-16 23/12/15 – 17/02/16	Cold 2016-17 01/11/16-18/03/17
BBOA	12.4%	8.9%	11.9%
HOA	12.2%	9.7%	16.4%
COA	10.4%	8.1%	11.7%
SV-OOA	19.8%	17.7%	28%
LV-OOA	45.2%	55.6%	32%