Supplemental Material
Aerosol distribution in the northern Gulf of Guinea: local anthropogenic
sources, long-range transport and the role coastal shallow circulations
Cyrille Flamant ¹ , Adrien Deroubaix ^{1,2} , Patrick Chazette ³ , Joel Brito ⁴ , Marco Gaetani ¹ ,
Peter Knippertz ⁵ , Andreas H. Fink ⁵ , Gaëlle de Coetlogon ¹ , Laurent Menut ² , Aurélie
Colomb ⁴ , Cyrielle Denjean ⁶ , Remi Meynadier ¹ , Philip Rosenberg ⁷ , Regis Dupuy ⁴ , Alfons
Schwarzenboeck₄ and Julien Totems³
¹ Laboratoire Atmosphères Milieux Observations Spatiales, Sorbonne Université,
Université Paris-Saclay and CNRS, Paris, France
² Laboratoire de Météorologie Dynamique, Ecole Polytechnique, IPSL Research
University, Ecole Normale Supérieure, Université Paris-Saclay, Sorbonne Université,
CNRS, Palaiseau, France
³ Laboratoire des Sciences du Climat et de l'Environnement, CEA, CNRS, Université
Paris-Saclay, Gif-sur-Yvette, France ⁴ Laboratoire de Météorologie Physique, Université
Clermont Auvergne, CNRS, Clermont-Ferrand, France
⁵ Institute of Meteorology and Climate Research, Karlsruhe Institute of Technology,
Karlsruhe, Germany
⁶ Centre National de Recherches Météorologiques, Météo-France and CNRS,
Toulouse, France
⁷ Institute of Climate and Atmospheric Science, School of Earth and Environment,
University of Leeds, Leeds, United Kingdom

27 Comparison between WRF dynamics at 2-km, aircraft and radiosounding
 28 observations

29



Figure S1: Evolution of temperature (a) and wind speed (b) along the ATR 42 flight track for in situ observations (closed black circles) and WRF (blue solid line). The variability of the observations within WRF grid cells is indicated as whiskers.





Figure S2: Wind speed (left column) and wind direction profiles (right column) in Accra (a and b, respectively) and Cotonou (c and d, respectively). Observations are shown as black dots and 2-km WRF simulations are shown as red solid lines. The radiosoundings were released at 1700 and 1612 in Accra and Cotonou, respectively.

- 39
- 40



(b)





- Figure S3: ECMWF IFS analyses at 1200 UTC on 2 July for water vapor mixing ratio (kg
 kg⁻¹, color shading), wind (vectors) and geopotential height (m, blue solid lines) at
 (a) 925 hPa and (b) 600 hPa. The African continent coastline is shown as a black solid
 line and cities of interest are indicated.



Figure S4: ECMWF CAMS forecast on 2 July 2016 at 1200 UTC (+12h forecast) for (a)
organic matter AOD and (b) dust AOD. The African continent coastline is shown as a
black solid line and cities of interest are indicated.

55 Regional overturning circulation induced by land-sea skin temperature gradients





Figure S5: Left column: West-east oriented vertical cross section (1000-500 hPa) of 58 59 zonal-vertical wind vectors from IFS analyses (blue) between 5°W and 10°E averaged 60 between 4.54°N and 6.28°E at (a) 0600 and (c) 1200 UTC on 2 July 2016. The thick red line is the projection of the ATR 42 aircraft track onto the cross-section. The thick 61 green and blue lines at the bottom of the graph indicate the presence of land and 62 ocean, respectively. Surface characteristics are defined based on the dominating 63 surface type in the latitudinal band considered for the average of the wind field. 64 65 Right column: IFS skin temperature (colors) and wind field at 10 m (vectors) at (b) 0600 UTC and (d) 1200 UTC. Cross-sections (a) and (c) are computed in the zonal box 66 shown in Figure 12c in the main paper. 67





Figure S6: Left column: Copernicus skin temperature at (b) 0600 UTC, (d) 1200 UTC and (f) 1800 UTC on 2 July 2016. Right column: IFS minus Copernicus skin temperature at (b) 0600 UTC, (d) 1200 UTC and (f) 1800 UTC. IFS skin temperature, originally at 0.125° resolution, has been linearly interpolated onto the Copernicus grid at 5 km before computing the differences.



77 Figure S7: MODIS-derived SST on 2 July 2016, with superimposed ATR 42 flight track

