

1 **The size resolved cloud condensation nuclei (CCN)**  
2 **activity and its prediction based on aerosol**  
3 **hygroscopicity and composition in the Pearl Delta River**  
4 **(PRD) Region during wintertime 2014**

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18  
19 **Abstract.** A hygroscopicity-tandem differential mobility analyzer (H-TDMA), a scanning mobility  
20 CCN analyzer (SMCA), and an aerodyne high resolution time-of-flight aerosol mass spectrometer  
21 (HR-ToF-AMS) were used to respectively measure the hygroscopicity, condensation nuclei activation,  
22 and chemical composition of aerosol particles at the Panyu site in the Pearl River Region during  
23 wintertime 2014. The distribution of the size-resolved cloud condensation nuclei (CCN) at four  
24 supersaturations (SS=0.1%, 0.2%, 0.4%, and 0.7%) and the aerosol particle size distribution were  
25 obtained by the SMCA. The hygroscopicity parameter  $\kappa$  ( $\kappa_{\text{CCN}}$ ,  $\kappa_{\text{H-TDMA}}$ , and  $\kappa_{\text{AMS}}$ ) was respectively

1 calculated based upon the SMCA, H-TDMA, and AMS measurements. The results showed that the  
2  $\kappa_{\text{H-TDMA}}$  value was slightly smaller than the  $\kappa_{\text{CCN}}$  one at all diameters and for particles larger than 100  
3 nm and the  $\kappa_{\text{AMS}}$  value was significantly smaller than the others ( $\kappa_{\text{CCN}}$  and  $\kappa_{\text{H-TDMA}}$ ), which could be  
4 attributed to the underestimated hygroscopicity of the organics ( $\kappa_{\text{org}}$ ). The activation ratio (AR)  
5 calculated from the growth factor – probability density function (Gf-PDF) without surface tension  
6 correction was found to be lower than that from the CCN measurements, due most likely to the  
7 uncorrected surface tension ( $\sigma_{\text{s/a}}$ ) that did not consider the surfactant effects of the organic compounds.  
8 We demonstrated that better agreement between the calculated and measured AR could be obtained by  
9 adjusting  $\sigma_{\text{s/a}}$ . Various schemes were proposed to predict the CCN number concentration ( $N_{\text{CCN}}$ ) based  
10 on the H-TDMA and AMS measurements. In general, the predicted  $N_{\text{CCN}}$  agreed reasonably well with  
11 the corresponding measured ones using different schemes. For the H-TDMA measurements, the  $N_{\text{CCN}}$   
12 value predicted from the real time AR measurements was slightly smaller (~6.8%) than that from the  
13 activation diameter ( $D_{50}$ ) method due to the assumed internal mixing in the  $D_{50}$  prediction. The  $N_{\text{CCN}}$   
14 values predicted from bulk chemical composition of  $\text{PM}_{10}$  were higher (~11.5%) than those from  
15 size-resolved composition measured by the AMS because a significant fraction of  $\text{PM}_{10}$  was composed  
16 of inorganic matter. The  $N_{\text{CCN}}$  values calculated from AMS measurement were under-predicted at 0.1%  
17 and 0.2% supersaturations, which could be due to underestimate of  $\kappa_{\text{org}}$  and overestimate of  $\sigma_{\text{s/a}}$ . For  
18  $\text{SS}=0.4\%$  and  $0.7\%$ , slight over-predicted  $N_{\text{CCN}}$  values were found because of the internal mixing  
19 assumption. Our results highlight the need for accurately evaluating the effects of organics on both the  
20 hygroscopic parameter  $\kappa$  and the surface tension  $\sigma$  in order to accurately predict CCN activity.  
21

## 1 1 Introduction

2 Aerosol particles can directly impact global climate by scattering and absorbing solar radiation  
3 (Stocker, 2013), while they can influence cloud formation, life time and optical properties by acting as  
4 cloud condensation nuclei (CCN), indirectly exerting climatic forcing on the Earth's atmosphere. In  
5 general, aerosol particles increase the CCN concentration and hence cause cooling effects on the global  
6 radiation balance. However, to what extent aerosol particles contribute to the radiation forcing is still  
7 highly uncertain (Stocker, 2013). It is hence important to measure chemical composition and properties  
8 of aerosol particles in order to assess their abilities of acting as CCN and contribution to cloud  
9 formation, further facilitating our understanding of the impacts of atmospheric aerosols on regional and  
10 global climate.

11

12 The extent to which aerosol particles can affect cloud formation is dependent on their fraction that can  
13 be activated to become CCN. This fraction of activation is termed as CCN activity that is determined  
14 by the chemical composition, sizes, and the water saturation ratio of the particles (Farmer et al., 2015).

15 The size-dependent saturation ratio (S) can be calculated from the Köhler equation (Köhler, 1936):

$$16 \quad S = a_w \exp\left(\frac{4\sigma_{s/a}M_w}{RT\rho_w D}\right) \quad (1)$$

17 where  $a_w$  is the water activity in solution,  $\sigma_{s/a}$  is surface tension of the solution/air interface,  $M_w$  is  
18 the mole weight of water, R is the universal gas constant, T is temperature in Kelvin, and D is the  
19 diameter of the droplet. The  $a_w$  represents Raoult effect, which means that the activation potential  
20 increases with the concentration of the solution. The term  $\exp\left(\frac{4\sigma_{s/a}M_w}{RT\rho_w D}\right)$  represents Kelvin effect,  
21 which relates the surface curvature to the saturation vapor pressure of the droplet. The activation

1 potential increases with increase of the droplet diameter or decrease of surface tension  $\sigma_{s/a}$  and the  
2  $\sigma_{s/a}$  value is sensitive to the organic surfactant effect. The two important parameters, the water activity  
3 ( $a_w$ ) and surface tension ( $\sigma_{s/a}$ ), are dependent on the composition of the aerosol particles, assuming  
4 those particles have the same properties as their corresponding bulk solutions. The effects of organics  
5 on the CCN activity have been extensively investigated; however, many outstanding questions still  
6 remain. Sorjamaa et al. (2004) suggested that the partitioning of surfactants had to be considered when  
7 evaluating the Kelvin effect and the Raoult effect. According to their experimental results, the  
8 surfactant partitioning could alter the Raoult effect and that the change is large enough to depress CCN  
9 activity. However, another experiment conducted by Engelhart et al. (2008) revealed that the organics  
10 in aged monoterpene aerosols could depress surface tension by about  $0.01 \text{ N m}^{-1}$  and hence increase  
11 CCN activity. Ovadnevaite et al. (2017) also presented observational and theoretical evidences that the  
12 decrease of surface tension could prevail over the Raoult effect, which led to the increase of CCN  
13 activity. Salma et al. (2006) isolated humic-like substances (HULIS) from PM<sub>2.5</sub> fraction aerosol  
14 samples and investigated the surface tension properties of the HULIS pure solutions. The results show  
15 that thermodynamic equilibrium on surface could only be reached after several hours. Because the  
16 depression of surface tension was controlled by diffusion of surfactants from the bulk of the droplet to  
17 its surface, the extent of the actual decrease of surface tension was hence kinetically limited. A hybrid  
18 model proposed by Petters and Kreidenweis (2013) was used to predict the effects of surfactants on the  
19 CCN activity. The model predicted strong effects of the surfactants on ternary systems where common  
20 ions were present. However, due to the limited measurement techniques, the available laboratory data  
21 were still not sufficient to support this prediction and more solid data were needed to validate the

1 surfactant effects on the CCN activity.

2

3 The CCN activity can be characterized by the hygroscopicity parameter  $\kappa$  that was initially proposed by  
4 Petters and Kreidenweis (2007). Aerosol hygroscopicity represents the ability of the particles to grow  
5 by absorbing water vapor from the atmosphere and the extent to which the particles are hygroscopic  
6 can be evaluated by the  $\kappa$  values, which can be determined from the H-TDMA or CCNc measurements.

7 The  $\kappa$  values were measured worldwide extensively either in the field measurements or in the  
8 laboratory experiments and depending on the organic content of the particles, a wide range of  $\kappa$  values  
9 were reported in the literatures. Cerully et al. (2011) showed that the  $\kappa$  values measured in 2007 by  
10 Flow-Streamwise Thermal Gradient CCN Chamber (CFSTGC) ranged mostly between 0.1 and 0.4 in a  
11 forest environment in Finland. Hong et al. (2014) obtained the average  $\kappa$  values of 0.15 (110 nm) and  
12 0.28 (102 nm) measured by H-TDMA at the same site in 2010. Chang et al. (2010) used an AMS to  
13 measure aerosol chemical composition and a mole ratio of atomic oxygen to atomic carbon (O/C) at a  
14 rural site in Canada. They reported a relationship between the  $\kappa$  values of organics and the O/C ratio as  
15  $\kappa_{\text{org}}=(0.29\pm 0.05)*(O/C)$ . Tritscher et al. (2011) conducted smog chamber experiments for  
16 measurements of the  $\kappa$  values of aging secondary organic aerosols and they found that the  $\kappa$  was a  
17 sensitive indicator of the SOA properties.

18

19 Although the  $\kappa$  values were reported under different environments in many locations, only a few  
20 studies were conducted to measure  $\kappa$  in the Pearl River Delta (PRD) region (Cheung et al. 2015;  
21 Schurman et al. 2017). Jiang et al. (2016) compared the  $\kappa$  values between wintertime (0.18-0.22) and

1 summertime (0.17-0.21) in Guangzhou. Cai et al. (2017) reported the  $\kappa$  values of about 0.4-0.6 and  
2 0.2-0.3 measured by the H-TDMA respectively in Cape Hedo (Japan) and in Guangzhou (China).  
3 Alternatively, the average  $\kappa$  values can be predicted by the ZSR mixing rule (Zdanovskii, 1948; Stokes  
4 and Robinson, 1966) which is based on the chemical composition of the aerosol particles from the  
5 AMS measurements. Liu et al. (2014) reported the  $\kappa$  values of 0.22 to 0.32 using the ZSR mixing rule,  
6 consistent with the values (0.25 to 0.34) based on the H-TDMA measurements.

7

8 Once the  $\kappa$  values were determined, they could then be employed to predict the CCN activity that was  
9 characterized by two important parameters: activation diameter ( $D_{50}$ ) and activation ratio (AR). Until  
10 now, the CCN activity (thus the above three parameters) can be determined using the following three  
11 methods:

12 (1) The combination of Cloud Condensation Nuclei counter (CCNc) and Scanning Mobility Particle  
13 Sizer (SMPS). The CCN number was measured by the CCNc at different supersaturation ratios (SS,  
14 typically 0.05% ~ 1%). Meanwhile, the  $D_{50}$  and size-resolved activation ratios could be measured by  
15 combining the CCNc with a differential mobility analyzer (DMA) and a condensation particle counter  
16 (CPC) (Moore et al., 2011; Deng et al., 2011), referred to as Scanning Mobility CCN Analysis (SMCA)  
17 based on measurements from a SMPS (DMA+CPC) and a CCNc. This method can measure the  
18 size-resolved CCN distributions at a high time resolution (Moore et al., 2010) and has been applied in  
19 lab experiments (Asaawuku et al., 2009) and field campaigns (Moore et al., 2008) to measure CCN  
20 activity.

21 (2) The ZSR method based on chemical composition measurements. The CCN concentrations were

1 inverted from the chemical composition and the size distribution of the aerosol particles measured  
2 respectively from the aerosol mass spectrometer (AMS) and SMPS (Moore et al., 2012; Meng et al.,  
3 2014). The  $\kappa$  was then calculated from the ZSR mixing rule. In general, the particles were assumed to  
4 be internally mixed, which might lead to a large uncertainty (up to 80%) in predicting  $N_{CCN}$  in some  
5 cases (Wang et al., 2010).

6 (3) The H-TDMA method. The size-resolved CCN distribution and activation ratios could be  
7 determined from hygroscopicity and size distribution measured using the H-TDMA (Good et al., 2010;  
8 Wu et al., 2013). The H-TDMA measured the distribution of hygroscopic growth factor (Gf) at a fixed  
9 relative humidity for a selected diameter of aerosol particles. Väisänen et al. (2016) reported that the  
10 measured  $N_{CCN}$  with the H-TDMA agreed well with that from in-cloud prediction, where the sample  
11 was collected from a tower approximately 224 m above the surrounding lake level. On the other hand,  
12 Chan (2008) attributed differences in  $\kappa$  from HTDMA and CCN measurements to sparingly soluble  
13 organics that did not easily deliquesce in the previous measurements.

14

15 The PRD region is one of the most economically invigorating regions in China. This region is  
16 subjected to severe air pollution due to intensive human activities and insufficient pollution control  
17 measures. High particle loading leads to both visibility degradation and large cooling effects due to  
18 decrease of solar radiation. During wintertime, high concentrations of fine particles also cause severe  
19 haze events that pose health risk on people at the regional scale. It is hence an ideal location to  
20 investigate the influence of local anthropogenic emissions on the particles properties. However, there is  
21 still lack of understanding on the relationship between the CCN activity and its controlling factors (e.g.,

1 chemical composition and hygroscopicity of aerosol particles), hindering policy-makers to propose  
2 effective measures for climate-related policy-making.

3

4 In this study, we used the SMCA, H-TDMA, and HR-ToF-AMS to respectively measure CCN activity,  
5 hygroscopicity, and chemical composition. We reported the relationship between CCN activity and  
6 hygroscopicity/chemical composition of aerosol particles in the PRD region, where only a few studies  
7 on such relationship were available in the literature. The measurements were performed during  
8 wintertime 2014 (November and December). The CCN properties were predicted based on the  
9 combined SMCA, H-TMDA and HR-ToF-AMS measurements. The methods employed to predict the  
10 CCN concentrations were evaluated and the impact of organics on CCN concentrations was discussed.

11

12 **2 Experiments and data analysis**

13 **2.1 Measurement site**

14 The field measurements were conducted at the Chinese Meteorological Administration (CMA)  
15 Atmospheric Watch Network (CAWNET) Station in Panyu, Guangzhou, China, during wintertime  
16 2014 (November and December). The Panyu Station is located at the center of the PRD region and at  
17 the top of Dazhengang Mountain (23°00'N, 113°21'E) with an altitude of about 150 m. No significant  
18 local emission sources were around the site. Detailed description of the measurement site and  
19 instruments (i.e., the HTDMA and the AMS) can be found elsewhere (Cai et al., 2017; Qin et al. 2017).

20

21 **2.2 Instrumentation**



1    **2.2.1    Aerosol hygroscopicity measurements**

2    Size-resolved aerosol hygroscopicity and particle number size distribution (PNSD) were measured by a  
3    H-TDMA which was developed by Tan et al. (2013). The hygroscopicity data were only available in  
4    November due to the failure of the H-TDMA during December. An aerosol sampling port equipped  
5    with a PM<sub>1.0</sub> impactor inlet was used during the measurement period. Ambient sampling flow first  
6    passed through a Nafion dryer (Model PD-70T-24ss, Perma Pure Inc., USA) to achieve a RH of <10%.  
7    We considered the particles to be dry when the RH values were less than 10%. The particles were  
8    subsequently charged by a neutralizer (Kr85, TSI Inc.) and size-selected by a differential mobility  
9    analyzer (DMA1, Model 3081L, TSI Inc.). The mono-disperse particles with a specific diameter (D<sub>0</sub>)  
10    were then introduced into a Nafion humidifier (Model PD-70T-24ss, Perma Pure Inc., USA) under a  
11    fixed RH of (90 ±0.44) %. Another differential mobility analyzer (DMA2, Model 3081L, TSI Inc.) and  
12    a condensation particle counter (CPC, Model 3772, TSI Inc.) were used to measure the number size  
13    distribution of the humidified particles (D<sub>p</sub>). Thus, growth factor (G<sub>f</sub>) of the particles can be  
14    calculated:

15    
$$G_f = \frac{D_p}{D_0} \tag{2}$$

16  
17    During the campaign, we selected five dry mobility diameters (40, 80, 110, 150, and 200 nm) for the  
18    H-TDMA measurements. The measurements were performed continuously except for regular  
19    calibration of the instrument. We used standard polystyrene latex spheres and ammonium sulfate to  
20    perform the DMA calibration to ensure the instrument to function normally.

21

## 1 2.2.2 Size-resolved CCN activity measurements

2 Size-resolved CCN spectra and activation ratios were measured with the SMCA initially proposed by  
3 Moore et al. (2010)). In this work, the SMCA consisted of a CCNc-100 (DMT Inc.), a differential  
4 mobility analyzer (DMA, Model 3081L, TSI Inc.) and a condensation particle counter (CPC, Model  
5 3787, TSI Inc.). In the SMCA system, the combined DMA and CPC were used as a scanning mobility  
6 particle sizer (SMPS) during the measurements. The dry particles after the Nafion dryer were  
7 neutralized by the Kr85 neutralizer and were subsequently classified by the DMA. The mono-disperse  
8 particles were split into two streams: one to the CPC for measurement of total particle number  
9 concentration ( $N_{CN}$ ) and another to the CCNc-100 for measurements of the CCN number concentration.  
10 The aerosol and CPC flow rate was both  $1.0 \text{ L min}^{-1}$  for the DMA and the CPC ( $0.5 \text{ L min}^{-1}$  makeup  
11 flow and  $0.5 \text{ L min}^{-1}$  sample flow), respectively. The CCNc-100 drew another aerosol flow rate of  $0.5 \text{ L}$   
12  $\text{min}^{-1}$ . The SMCA was protocolled to measure particles at a mobility diameter range of 10 - 400 nm. The  
13 supersaturation in the CCNc-100 was set to be 0.1%, 0.2%, 0.4% and 0.7% respectively for each  
14 measurement cycle. The CCNc-100 was regularly calibrated with ammonium sulfate particles at the  
15 four SS (0.1%, 0.2%, 0.4%, and 0.7%). Previous studied showed that different parameterizations in the  
16 Köhler theory can retrieve different critical supersaturations (Rose et al., 2008; Wang et al., 2017).  
17 When performing the CCNc calibration, we assumed the density and molecular weight of ammonium  
18 sulfate to be  $1770 \text{ kg m}^{-3}$  and  $0.132141 \text{ kg mol}^{-1}$ , respectively. In the CCNc calibration, the water  
19 activity ( $a_w$ ) was approximated according to Rose et al. (2008),

$$20 \quad a_w = \exp(-i_s \mu_s M_w) \quad (3)$$

21 where  $i_s$ ,  $\mu_s$  and  $M_w$  is the van't Hoff factor, molality of solute, and the molar mass of water  
22 ( $0.01802 \text{ kg mol}^{-1}$ ), respectively. The van't Hoff factor  $i_s$  is calculated from a polynomial fit to Pitzer

1 model output data (Morre et al., 2010). In this study, we adopted the simplest parameterization of the  
2 surface tension of the solution (Rose et al., 2008), that is, it was simply approximated by the surface  
3 tension of pure water ( $0.072 \text{ N m}^{-1}$ ) according to Seinfeld and Pandis (2007).

4

5 We also set the temperature and the pressure to 298.15 K and 1026 hPa, respectively. A temperature  
6 gradient  $\Delta T$  of about 3-8 K in the CCNc column was also used in the calibrations. Similarly, the DMA  
7 was calibrated with standard polystyrene latex spheres before and after the campaign for quality  
8 assurance and control.

9

### 10 **2.2.3 Aerosol chemical composition measurements**

11 An Aerodyne high-resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS) was employed  
12 in the campaign to measure non-refractory  $\text{PM}_{10}$  chemical composition (bulk and size-resolved)  
13 including sulfate, nitrate, ammonium, chloride, and organics. The refractory components such as black  
14 carbon, sea salts and crustal species cannot be measured by this instrument. Detailed description of the  
15 HR-ToF-AMS can be found elsewhere (DeCarlo et al., 2006; Jimenez et al., 2003). Here only a brief  
16 description relevant to the measurements is given. The instrument was operated in three modes (pToF,  
17 V, and W mode). Particle size distribution could be obtained based on time-of-flight of the particles in  
18 pToF mode. In V and W modes, the resolving power of the mass spectrometer was approximately 2000  
19 and 4000, respectively. The instrument collected alternatively 5-min average mass spectra for the V +  
20 pToF modes and the W mode. The monodisperse pure ammonium nitrate ( $\text{NH}_4\text{NO}_3$ ) particles selected  
21 by a DMA (400 nm) were used weekly in the ionization efficiency (IE) calibration. Background signals

1 were obtained daily for about 30 minutes by introducing filtered ambient air with a HEPA filter in the  
2 sample flow. Before and after the measurement, the sampling flow rate was calibration with a Gilian  
3 gibrator. We also generated PSL (Duke Scientific) and ammonium nitrate particles in a size range of  
4 178~800 nm to calibrate the pToF size. Note that the mass concentrations were too low for particle  
5 diameters smaller than 65 nm and data for those particles were hence discarded in this study. A more  
6 detailed description of the AMS performance during the measurements can be found in Qin et al. (2017)  
7 and Cai et al. (2017).

8

9 The AMS measured size-resolved chemical composition of particles in vacuum aerodynamic diameter  
10 ( $D_{va}$ ). It is hence necessary to convert aerodynamic diameter to mobility diameter in order to compare  
11 the AMS data and the SMCA data. We adopted the equation derived by DeCarlo et al. (2004) to do the  
12 conversion. Here we assume a density of  $1700 \text{ kg m}^{-3}$  for particles measured by the AMS (DeCarlo et  
13 al., 2004).

14

## 15 **2.3 Data processing and methodology**

### 16 **2.3.1 Hygroscopicity**

17 Due to the effects of diffusing transfer function, the measured distribution function (MDF) given by  
18 H-TDMA is only a skewed and smoothed integral transform of the actual growth factor probability  
19 density function (Gf-PDF) of the particles (Gysel et al., 2009). Here the TDMAfit algorithm  
20 (Stolzenburg and McMurry, 2008) was applied to narrow the uncertainties caused by the diffusion  
21 broadening. The TMDAfit algorithm describes the Gf-PDF as a combination of several (usually smaller

1 than three) lognormal distribution functions, in which the parameters of each mode are considered as  
2 mean  $Gf$ , standard deviation, and number fraction. The detailed data inversion process of the H-TDMA  
3 instrument can be found in Tan et al. (2013). Note that we include multiply charged correction for the  
4 SMCA, SMPS and H-TDMA data when the data were inverted so that the contributions of the multiply  
5 charged particles were accounted for all the measured particle data.

6

7 As mentioned in the introduction, the CCN activity can be represented by a widely used hygroscopicity  
8 parameter  $\kappa$  (Petters and Kreidenweis, 2007). According to the  $\kappa$ -Köhler theory, for a known  
9 temperature,  $\kappa$  and  $Gf$  can be related via eq. 4 (Petters and Kreidenweis, 2007):

$$10 \quad \kappa = (Gf^3 - 1) \left[ \frac{1}{RH} \exp\left(\frac{4\sigma_{s/a}M_w}{RT\rho_w D} - 1\right) \right] \quad (4)$$

11 where  $\rho_w$  is the density of water (about 997.04 kg m<sup>-3</sup> at 298.15 K),  $M_w$  is the molecular weight of  
12 water (0.018 kg mol<sup>-1</sup>),  $\sigma_{s/a}$  is the surface tension of the solution/air interface and here pure water is  
13 tentatively assumed for the solution ( $\sigma_{s/a}$  = 0.0728 N m<sup>-1</sup> at 298.15 K),  $R$  is the universal gas constant  
14 (about 8.31 J mol<sup>-1</sup>K<sup>-1</sup>),  $T$  is thermodynamic temperature in Kelvin (298.15 K), and  $D$  is the particle  
15 diameter (in meter).

16

### 17 2.3.2 CCN activation

18 The  $N_{CN}$  and  $N_{CCN}$  data were respectively measured by the SMPS and the CCNc-100 and they were  
19 used to calculate the size-resolved CCN activation ratios (AR) which was defined as the ratio of  $N_{CCN}$   
20 to  $N_{CN}$  at each particle size. The activation ratio can be obtained by fitting the ratio with the sigmoidal  
21 function with respect to  $D_p$ :

1 
$$\frac{N_{CCN}}{N_{CN}} = \frac{B}{1 + \left(\frac{D_p}{D_{50}}\right)^C} \quad (5)$$

2 where  $D_p$  is the particle dry diameter,  $B$ ,  $C$  and  $D_{50}$  are fitting coefficients that represent the  
 3 asymptote, the slope, and the inflection point of the sigmoid, respectively (Moore et al., 2010). **Note**  
 4 **that the parameter  $C$  is a fitting coefficient with no specific physical meaning. However, a small  $C$**   
 5 **value indicates a steep activation curve.**  $D_{50}$  is also called the critical diameter or the activation  
 6 diameter, that is, the diameter at which 50% of the particles are activated at a specific SS.

7

8 Alternatively, the hygroscopicity parameter  $\kappa$  can be calculated from the critical saturation ratio ( $Sc$ )  
 9 and  $D_{50}$  from the following equation (Petters and Kreidenweis, 2007):

10 
$$\kappa = \frac{4A^3}{27D_{50}^3(\ln Sc)^2}, \quad A = \frac{4\sigma_s/aM_w}{RT\rho_w} \quad (6)$$

11

### 12 **2.3.3 CCN prediction based on H-TDMA and AMS measurements**

13 The  $N_{CCN}$  can be predicted based on either the aerosol hygroscopicity data (measured by the H-TDMA)  
 14 or the AMS data. Figure 1 is the schematic diagram of the four approaches we followed to predict  $N_{CCN}$   
 15 based on the above two measured datasets. In the first approach (I in Fig. 1), the mixing state and size  
 16 dependence were taken into account. We assumed the critical hygroscopicity parameter  $\kappa_{critical}$  to be a  
 17 function of the particle diameter and the supersaturation ratio (denoted as  $\kappa_{critical}(D_p, SS)$ ). The  $\kappa_{critical}$   
 18 was hence defined as the point at which all the particles were activated at a specific diameter and a  
 19 specific SS. Here we measured hygroscopicity using the H-TDMA at five dry diameters and the CCN  
 20 concentrations at four SS. We calculated the  $\kappa_{critical}(D_p, SS)$  using eq.6 for a known diameter and SS. A  
 21 particle with a  $\kappa$  value higher than  $\kappa_{critical}(D_p, SS)$  was considered to be activated as an CCN (Fig. 1a)

1 and the shadow area represented the particles which can be activated as CCN for a known diameter and  
 2 SS. The activation ratio for a specific diameter at a specific SS was obtained by integrating the  $\kappa$ -PDF  
 3 for  $\kappa > \kappa_{\text{critical}}(Dp, SS)$ .

4

5 This approach is similar to the one employed in Kammermann et al. (2010), however, we used the  
 6 size-resolved activation ratio ( $AR_{SR}$ ) to calculate the  $N_{CCN}$ . The  $AR_{SR}$  was determined by fitting the  
 7  $AR(Dp, SS)$  to the diameter  $Dp$  using eq. 5 for the five measured diameters (Fig. 1d). Thus, the  
 8 calculated  $N_{CCN}$  using the activation ratio can be expressed as (Fig. 1e):

$$9 \quad N_{CCN}(SS) = \int_0^{\infty} AR_{SR}(Dp, SS) N_{CN}(Dp) dDp \quad (7)$$

10

11 In the second approach (II in Fig.1), the particles were assumed to be internally mixed. The  $D_{50}$  was  
 12 determined by fitting the  $AR(Dp, SS)$  to the diameter  $Dp$  (Fig. 1d). The  $N_{CCN}$  was obtained by  
 13 integrating the cloud nuclei concentration for particles larger than  $D_{50}$  based on the particle size  
 14 distribution (Fig. 1f), according to the following equation (eq. 8):

$$15 \quad N_{CCN}(SS) = \int_{D_{50}}^{\infty} N_{CN}(Dp) dDp \quad (8)$$

16

17 In the third and fourth approaches (III and IV in Fig. 1), the particles were also assumed to be internally  
 18 mixed. We then calculated the  $\kappa$  value according to the ZRS rule (eq. 9) based on the AMS  
 19 measurements.

$$20 \quad \kappa = \sum_i \varepsilon_i \kappa_i \quad (9)$$

21 where  $\varepsilon_i$  is the volume fraction of each component in the particles,  $\kappa_i$  is the  $\kappa$  value of each component.

1

2 The AMS only provided the ion concentrations during the measurements, while the ZSR rule required  
3 the volume fraction and hygroscopicity of each component. A simplified ion pairing scheme developed  
4 by Gysel et al. (2007) was used to reconstruct the  $\text{NH}_4^+$ ,  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  measured by the AMS:

$$n_{\text{NH}_4\text{NO}_3} = n_{\text{NO}_3^-}$$

$$n_{\text{H}_2\text{SO}_4} = \max(0, n_{\text{SO}_4^{2-}} - n_{\text{NH}_4^+} + n_{\text{NO}_3^-})$$

$$n_{\text{NH}_4\text{HSO}_4} = \min(2n_{\text{SO}_4^{2-}} - n_{\text{NH}_4^+} + n_{\text{NO}_3^-}, n_{\text{NH}_4^+} - n_{\text{NO}_3^-})$$

$$n_{(\text{NH}_4)_2\text{SO}_4} = \max(n_{\text{NH}_4^+} - n_{\text{NO}_3^-} - n_{\text{SO}_4^{2-}}, 0)$$

5  $n_{\text{HNO}_3} = 0,$  (10)

6 where  $n$  denotes the number of moles of each component (i.e.,  $\text{NH}_4^+$ ,  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$ ). Here we used  
7 the ADDEM proposed by Topping et al. (2005) to calculate the  $\kappa$  values of the inorganic species and  
8 those of the organics were tentatively assumed to be 0.1 (Meng et al., 2014). Table 1 lists the  $\kappa$  values  
9 of the relevant species used in the study based on the calculations and the above assumption.

10

11 Here instead of being determined from fitting of  $\text{AR}_{\text{SR}}$  to  $D_p$  used in the second approach, the  $D_{50}$  was  
12 calculated from the above  $\kappa$  values using eq. 6. In the third approach, the  $\kappa$  values were size-resolved  
13 because the chemical composition of the particles was size dependent (Fig. 1b). In the fourth approach,  
14 the particles were assumed to have the same chemical composition and hygroscopicity as those in  $\text{PM}_{10}$   
15 (Fig. 1c). The  $N_{\text{CCN}}$  was then predicted using eq. 8 (Figs. 1g and h).

16

17 **3 Results and discussion**



### 1 3.1 Overview

2 Table 2 summarizes the observed CCN activity during the campaign. Overall, the average  $N_{\text{CCN}}$  at 0.1,  
3 0.2, 0.4, and 0.7% SS were about 3100, 5100, 6500, and 7900  $\text{cm}^{-3}$ , respectively. The average  
4 activation ratios (AR) at the above four SS were 0.26, 0.41, 0.53 and 0.64, respectively. The average  
5  $D_{50}$  at the above four SS were 156, 107, 78 and 58 nm, respectively. The  $N_{\text{CCN}}$  at 0.7% SS was  
6 respectively lower than those of the previous measurements (10731  $\text{cm}^{-3}$  at 0.67% SS) in July 2006 in  
7 Guangzhou (Rose et al., 2010), but much higher than those measured (2085  $\text{cm}^{-3}$  at 0.7% SS) in May  
8 2011 in Hong Kong (Meng et al., 2014), while the  $N_{\text{CCN}}/N_{\text{CN,tot}}$  was lower than those from the previous  
9 measurements (at 0.67% SS) in Guangzhou (0.59, Rose et al., 2010) and similar to those from the  
10 measurements (at 0.7% SS) in Hong Kong (0.64, Meng et al., 2014). The  $D_{50}$  was larger than that in the  
11 previous measurements in Guangzhou (49 nm) and in Hong Kong (47 nm), due to the lower particle  
12 hygroscopicity in Guangzhou. The differences of the  $\kappa_{\text{CCN}}$  values between the two measurements (0.21  
13 in this winter campaign vs 0.28 during the summer season in Guangzhou both at 0.7% SS) suggested  
14 that the particles in the summer were in general more hygroscopic and hence were more readily  
15 activated than those in the winter, implying different chemical composition of the particles between the  
16 two distinct seasons. Mochida et al. (2010) measured the size-resolved CCN activity in Cape Hedo, a  
17 remote marine site rarely affected by anthropogenic emissions. The results showed that the  $D_{50}$  at 0.1%  
18 SS in Cape Hedo was about 130 nm, much larger than that in Guangzhou, leading to higher  
19 hygroscopicity of atmospheric particles in Cape Hedo than that in Guangzhou.

20

21 Figure 2 shows the average mass fraction of NR-PM<sub>1</sub> bulk composition and size-resolved (64-731 nm)

1 composition. The organics was dominant in the bulk NR-PM<sub>1</sub> (50%), followed by sulfate (26%) and  
2 nitrate (12%) (Fig. 2a). The mass fraction of the organics decreased with the size (Fig. 2b), from 73%  
3 at 64 nm to 42% at 397 nm. The mass fraction of organics at 397 nm was close to that of NR-PM<sub>1</sub> bulk,  
4 due to the fact that the PM<sub>1</sub> mass is dominated by particles in a diameter range of 200~500 nm (Tan et  
5 al., 2016). In comparison, the dominant NR-PM<sub>1</sub> species observed in Hong Kong were sulfate (51.0%)  
6 and organics (28.2%) (Lee et al., 2013), significantly different from our measurements, due probably to  
7 different origins of the dominant air masses between the two seasons. The measurement site in  
8 Guangzhou was impacted predominantly by the air mass from north, where straw burning contributes  
9 to a high mass fraction of organics matter (Cao et al., 2008).

10

11 Figure 3 shows the  $\kappa$  values based respectively on the CCN ( $\kappa_{\text{CCN}}$ ), AMS ( $\kappa_{\text{AMS}}$ ), and H-TDMA  
12 ( $\kappa_{\text{HTDMA}}$ ) measurements, along with the measured particle number size distribution (PNSD, 10-400 nm)  
13 during the campaign. The shadow area represents the interquartile range of the PNSD. A distinct peak  
14 at around 90 nm was observed from the PNSD (Fig. 3). The  $\kappa_{\text{AMS}}$  was calculated based on the  
15 size-resolved chemical composition, assuming the particles are internally mixed. At 0.7% SS, the D<sub>50</sub>  
16 was about 58 nm. Hence no  $\kappa_{\text{AMS}}$  was reported at this SS since we only measured particle composition  
17 above 63 nm using the AMS in this study. The  $\kappa$  values were shown in the interquartile range, with the  
18 largest variation from the CCN measurements (Fig. 3). As shown in Fig. 3, the difference between  $\kappa_{\text{CCN}}$   
19 and  $\kappa_{\text{H-TDMA}}$  is statistically insignificant at all employed diameters, while the one between  $\kappa_{\text{AMS}}$  and  
20  $\kappa_{\text{CCN}}$  became statistically significant at larger sizes of the particles. **The p-value between the  $\kappa_{\text{AMS}}$  and  
21 the  $\kappa_{\text{ccn}}$  was close to 0, indicating that the correlation between them is significant.** The  $\kappa_{\text{HTDMA}}$  values

1 were lower than those of the corresponding  $\kappa_{\text{CCN}}$  at most of the particle sizes, consistent with the  
2 previous observation (Pajunoja et al., 2015). This was probably due to the facts that particles contain a  
3 certain fraction of low solubility composition, such as secondary organic aerosols (SOA), contributing  
4 differently to hygroscopic growth and CCN activation. The available AMS data (Fig. 3) show that the  
5  $\kappa_{\text{AMS}}$  values were lower than the corresponding  $\kappa_{\text{CCN}}$  and  $\kappa_{\text{HTDMA}}$  values at all size ranges and the  
6 differences become larger with increasing particle sizes. This was probably due to underestimated  
7 hygroscopicity in the organic composition when using the AMS data, since we assumed a  $\kappa$  value of  
8 0.1 for all organics at all particle sizes. The hygroscopicity increased with particle diameters due to  
9 aerosol aging which increased the hygroscopic organic contents. Previous studies showed that the  $\kappa_{\text{org}}$   
10 values of larger particles are lower than those for smaller particles (Lance et al., 2013; Zhao et al. 2015)  
11 and hygroscopicity of organics is often found to be related to its chemical composition (f44 or O/C) in  
12 both field and laboratory studies (Chang et al. 2010; Massoli et al., 2010; Lambe et al., 2011; Mei et al.,  
13 2013, and others reference therein). We showed that the f44 increased with the particle size from the  
14 AMS data (Fig. S1). Note that the f44 for particle diameters smaller than 100 nm was discarded due to  
15 the data quality. The results indicate that the degree of oxidation of the organics was higher for larger  
16 size particles and the hygroscopicity for larger particles is higher (Chang et al., 2010). The measured  
17  $\kappa_{\text{mean}}$  values fall in a range of 0.22-0.30 for the particle sizes of 40-200 nm measured by H-TDMA in  
18 this study. The other aerosol hygroscopicity measurement in PRD (Jiang et al., 2016) reported the  $\kappa_{\text{mean}}$   
19 values ranging from 0.18 to 0.22 in 2012 winter season and 0.17 to 0.21 in 2013 summer season,  
20 suggesting an increase of the aerosol hygroscopicity, which might result from an increasing mass  
21 fraction of nitrate in recent years (Zhang et al., 2015; Itahashi et al., 2018), although the fraction

1 decrease of less hygroscopic compounds is not as significant as the fraction increase of the nitrate.  
2 However, the fraction of the non-hygroscopic compounds (i.e. EC) decreases more rapidly than the  
3 organic compounds.

4

5 Figure 4 shows the activation ratios (AR) measured by SMCA at four supersaturation ratios (0.1, 0.2,  
6 0.4, 0.7%) for particles below 300 nm. The activation curves obtained in this study were segmented  
7 into three sections: a steady rise at low ARs, a middle sharp increase, and a plateau at almost 100% AR.

8 We defined the steepness as the rate at which the AR increased with the particle sizes. Figure 4 shows  
9 the steepness increased with the SS, indicating that the curves became steeper with the SS and a larger

10 variation of the  $D_{50}$  was expected. In addition, the CCN activity was more sensitive to particle  
11 diameters at higher SS, which can be seen from partial derivative of  $\kappa_{critical}$  by  $\partial D_{50}$  (eq. 11):

$$12 \quad \frac{\partial \kappa_{critical}}{\partial D_{50}} = -\frac{4A^3}{9D_{50}^4 (\ln Sc)^2} \quad (11)$$

13

14 For a certain SS, the  $\kappa_{critical}$  value became more sensitive to  $D_{50}$  with decrease of the  $D_{50}$ .

15 Meanwhile, a high SS usually led to a low  $D_{50}$ . Therefore, the AR would vary with  $D_p$  more readily at  
16 higher SS and the curve would become steeper. A higher SS allowed a smaller particle to be activated  
17 and the activation curve became steeper, and vice versa for a lower SS.

18

19 The steepness of activation curve was also associated with the heterogeneity of aerosol chemical  
20 composition, that was, a steeper activation curve meant that aerosol particles had higher similarity in  
21 hygroscopicity. A bimodal distribution (peaks at about 1-1.1, and 1.5-1.7 Gf) of the Gf-PDFs was

1 observed along the  $Gf$  coordinate at all the five sizes of the particles measured by H-TDMA in this  
 2 study (Fig. 5), corresponding respectively to the less- and more-hygroscopic modes. Larger size  
 3 particles contain higher fractions of more-hygroscopic inorganics matters which lead to the increase of  
 4  $Gf$  of more-hygroscopic mode. The less-hygroscopic mode usually represents externally mixed black  
 5 carbon or fresh organics. Thus the less-hygroscopic mode for larger size particles more likely  
 6 represents the external mixing non-hygroscopic black carbon with a  $Gf$  value of 0.8-1.1, indicating that  
 7 the  $Gf$  of less-hygroscopic mode decreased and that of more-hygroscopic mode increased with the  
 8 particle diameter (Fig. 5). Since less-hygroscopic particles were usually associated with externally  
 9 mixed black carbon (BC) or fresh organics and more-hygroscopic particles usually represent the  
 10 inorganics matters or BC coated with inorganics matters (internally mixed). The decrease of peak area  
 11 of less-hygroscopic mode and the increase of more-hygroscopic mode indicate that the number fraction  
 12 of less-hygroscopic particles decreased while the more-hygroscopic particles fraction increased. Thus,  
 13 the particles became more internally mixed. Here a parameter  $\sigma$  is introduced to illustrate the deviation  
 14 of  $Gf$ -PDF (Gysel et al., 2009):

$$15 \quad Gf_{mean} = \int_0^{\infty} Gf c(Gf) dGf \quad (12-1)$$

$$16 \quad \sigma = \left( \int_0^{\infty} (Gf - Gf_{mean})^2 c(Gf) dGf \right)^{\frac{1}{2}} \quad (12-2)$$

17 where the  $c(Gf)$  denotes  $Gf$ -PDF and  $Gf_{mean}$  denotes number weighted mean  $Gf$ . The  $\sigma$  was  
 18 employed as a measure of the spread of  $Gf$ -PDF which represents the heterogeneity of aerosol chemical  
 19 composition (Sjogren et al., 2008; Liu et al., 2011). A small  $\sigma$  indicated that the heterogeneity of  
 20 aerosol chemical composition was low and aerosol particles had higher similarity in hygroscopicity.  
 21 The parameter  $C$  determined the shape of the activation curve which was segmented into steep and

1 smooth parts. A small  $C$  value means a steep activation curve and vice versa. Here an activation curve  
2 was assumed to be steep when the  $C$  values are lower than the lower quartile of all the  $C$  values,  
3 while the activation curve was considered to be smooth when the  $C$  values are higher than the upper  
4 quartile of all the  $C$  values. Table 3 summarizes the  $\sigma$  values of Gf-PDF for the corresponding steep  
5 and smooth activation curve at the four supersaturations. In general, the  $\sigma$  increased with the diameter,  
6 indicating that larger particles had higher heterogeneity of aerosol chemical composition. Meanwhile,  
7 the  $\sigma$  values for smooth curve were generally higher than the  $\sigma$  values for steep curve. The results  
8 implied that the shapes of activation curves were related to the heterogeneity of aerosol chemical  
9 composition. Cai et al. (2017) compared the Gf-PDF between Guangzhou and Cape Hedo and the  
10 results showed that only more-hygroscopic (MH) particles were observed in Cape Hedo, indicating that  
11 atmospheric particles tend to be more internally mixed in Cape Hedo than in Guangzhou. Meanwhile,  
12 atmospheric particles in Guangzhou have a higher degree of external mixing affected by more  
13 anthropogenic emissions, which in turn affect the CCN activity.

14

### 15 **3.2 Impact of organics on CCN activity**

16 Figure 6 shows the relationship between the  $D_{50}$  obtained from the SMCA measurements and the  
17 size-resolved mass fractions of organics ( $f_{\text{org}}$ ) at three supersaturation ratios (0.1%, 0.2%, and 0.4% SS).  
18 In general, the  $D_{50}$  increased with  $f_{\text{org}}$  at the three SS, with a slope of 127, 66, 21, and a fitting  
19 coefficient ( $R^2$ ) of 0.47, 0.31, 0.1 at 0.1%, 0.2%, and 0.4% SS, respectively. The particles usually  
20 became less hygroscopic with increase of the organic fractions ( $f_{\text{org}}$ ), which then required larger  
21 particles to be activated. At lower SS, better correlations were found between the  $f_{\text{org}}$  and the  $D_{50}$

1 because the  $D_{50}$  was more sensitive to hygroscopicity (The activation ratios increase more slowly with  
2 particle sizes at lower SS as shown in Fig. 4). It was hence more obvious at lower SS that the  
3 modification of the particle hygroscopicity caused by the change of the mass fraction of organics  
4 matter could greatly modify the  $D_{50}$  which might further affect the CCN activity. At higher SS,  
5 according to eq. 6, particles were more easily activated as CCN and the change of particles  
6 hygroscopicity would not significantly alter the CCN activity.

7

8 Organics can affect the CCN activity via two opposite ways: they can decrease the CCN activity by  
9 increasing the less hygroscopic organic fraction of the particles and thus increase the  $D_{50}$  as shown in  
10 Fig. 6; they can also increase the CCN activity by decreasing the surface tension of the particles. The  
11 latter effect has been demonstrated experimentally. For example, an increase of CCN activity was  
12 observed when organics were added to sulfate ammonium (Engelhart et al., 2008). In this study, we  
13 investigated the impacts of organics on CCN activity through adjusting the value of surface tension  
14 until the calculated AR values based on H-TDMA measurements agree with those obtained from  
15 SMCA measurements (measured AR). The calculated AR values were systematically lower than the  
16 corresponding measured ones because here the surface tension of bulk pure water ( $0.072 \text{ N m}^{-1}$ ) was  
17 assumed when calculating the AR from the H-TDMA measurements (Fig. 7). Note that the surface  
18 tension is not the only factor that determines the AR and other factors such as the sparingly soluble  
19 compounds in the particles may contribute to the AR, although they are currently not understood.  
20 Previous studies found that the hygroscopicity of the particles measured by the H-TDMA could be  
21 lower than that measured by the CCNc (Chan et al., 2008; Pajunoja et al. 2015; Petters et al., 2009;

1 Hansen et al., 2015; Hong et al., 2014) which might be attributed to low soluble compounds in the  
 2 particles. The deviation of the calculated AR from the measured AR is probably dependent on the  
 3 degree of dissolution of particles and the oxidative state of the organics in the particles.

4

5 The surface tension of a nanoparticle was substantially different from that of its bulk solution due to the  
 6 curvature effect (Ahn et al., 2010; Bogdan, 1997). The effects of size and composition on the surface  
 7 tension were currently not well understood. Here we proposed an approach to evaluate the impact of  
 8 organics on the surface tension ( $\sigma_{s/a}$ ) based on the fraction change of the calculated AR to the measured  
 9 AR. We defined this fraction change ( $\delta_{AR}$ ) as a function of surface tension, diameter, and  
 10 supersaturation:

$$11 \quad \delta_{AR}(\sigma_{s/a}, Dp, SS) = \frac{AR_m(Dp, SS) - AR_c(\sigma_{s/a}, Dp, SS)}{AR_m(Dp, SS)} \times 100\% \quad (13)$$

12 where  $AR_m(Dp, SS)$  is the measured AR for a certain diameter and SS,  $AR_c(\sigma_{s/a}, Dp, SS)$  is the  
 13 calculated AR for a certain diameter, SS, and  $\sigma_{s/a}$ . We excluded particles at the size of 200 nm because  
 14 they were easily activated even at 0.1% SS and the  $\delta_{AR}$  was expected to be independent of  $\sigma_{s/a}$ . Here the  
 15  $\sigma_{s/a}$  value varied between 0.03 and 0.072 N m<sup>-1</sup> (surface tension of pure water). Figure 8 shows the  $\delta_{AR}$   
 16 as a function of  $\sigma_{s/a}$  for the four particle diameters (40, 80, 110, 150 nm). The  $\delta_{AR}$  decreased with  
 17 increase of the  $\sigma_{s/a}$  for all given particle sizes, changing more rapidly for smaller particles (i.e., from  
 18 200% to -100% for 40 nm) than bigger particles (i.e., from 20% to -10% for 150 nm). The R<sup>2</sup> between  
 19 measured AR and predicted AR for a certain diameter and four supersaturations at  $\sigma_{s/a}=0.072$  N m<sup>-1</sup>  
 20 were 0.35, 0.93, 0.95 and 0.91, respectively. The  $\delta_{AR}$  values reached zero when the  $\sigma_{s/a}$  was set to be  
 21 about 0.054 N m<sup>-1</sup> for 40, 80, and 110 nm particles, and 0.062 N m<sup>-1</sup> for 150 nm particles, with a R<sup>2</sup> of



1 0.88, 0.94, 0.94 and 0.88 respectively. As a compromise, here we adopt a  $\sigma_{s/a}$  value of  $0.058 \text{ N m}^{-1}$   
2 (denoted as  $\sigma_{s/a}^*$ ) to predict AR. This  $\sigma_{s/a}^*$  value increased significantly the  $R^2$  compared to that based  
3 on pure water assumption ( $0.072 \text{ N m}^{-1}$ ) for 40 nm particles, while it was reasonable well for other  
4 sizes of particles (80, 110, 150 nm). The AR was then recalculated using the  $\sigma_{s/a}^*$  value and the  
5 prediction was significantly improved (Fig. 9). We calculated the p-value between the measured and  
6 predicted AR and the results showed that the p-value is close to 0, indicating a significant correlation  
7 between the two variables. The results demonstrated that partitioning of organics into aerosol particles  
8 would decrease their surface tension. Therefore, the pure water assumption for surface tension would  
9 lead to high uncertainties when it applied to predict the activation ratios of the aerosol particles at a  
10 certain size. Note that we did not consider the effects of individual organics due to the limited data  
11 from the chemical composition measurements. How chemical composition affects the surface tension  
12 of the particles is yet to be investigated.

13

### 14 **3.3 The $N_{CCN}$ prediction**

#### 15 **3.3.1 The $N_{CCN}$ prediction based on the H-TDMA measurements**

16 In this study, we used several approaches to predict the  $N_{CCN}$  based on the H-TDMA measurements,  
17 from either the activation curve or the  $D_{50}$ . Table 4 summarizes the methods that were used to predict  
18 the  $N_{CCN}$ , along with the slope and  $R^2$  between the predicted and the measured values. The mixing state  
19 of the aerosol particles is an important parameter in determining the  $N_{CCN}$ . The prediction of  $N_{CCN}$   
20 using activation curve means the  $N_{CCN}$  was calculated based on eq. 7. Meanwhile the prediction of  
21  $N_{CCN}$  using the  $D_{50}$  means that the  $N_{CCN}$  was calculated based on eq. 8 and the  $D_{50}$  was determined from

1 fitting the size-resolved activation ratio by eq. 5. The activation curve represented actual mixing state,  
2 while the  $D_{50}$  approach assumed that all particles were internally mixed. Scheme 5 in Table 4 was the  
3 method based on the activation curve with the new  $\sigma_{s/a}^*$  ( $0.058 \text{ N m}^{-1}$ ). Eq. 7 and Eq. 8 were  
4 respectively used to calculate the  $N_{CCN}$  following schemes 1, 2, 5, and the rest of the schemes. Scheme  
5 5 (real time activation curve using  $\sigma_{s/a}^*$ ) provided the best  $N_{CCN}$  predicted value (closest to the  
6 measured one), followed by scheme 3 (real time  $D_{50}$ ) > scheme 4 (average  $D_{50}$ ) > scheme 1 (real time  
7 activation curve) > scheme 2 (average activation curve). The  $R^2$  values for all the approaches were in  
8 general high (around 0.93). The CCN prediction based on scheme 2 led to the largest underestimation  
9 over the measured values. In general, the real time data (schemes 1 and 3) gave better predicted  $N_{CCN}$   
10 than the corresponding average data (schemes 2 and 4).

11

12 Figure 10 shows the correlation between the measured  $N_{CCN}$  and the predicted  $N_{CCN}$  from schemes 1-5  
13 at the four SS. For schemes 1-4, the predicted  $N_{CCN}$  values were found to be the largest deviation from  
14 the corresponding measured ones at 0.1% SS among all the approaches, probably due to the pure water  
15 assumption for surface tension ( $\sigma_{s/a} = 0.072 \text{ N m}^{-1}$ ). Meanwhile, because the CCN activity was sensitive  
16 to hygroscopicity of the particles at low SS, the uncertainties of hygroscopicity data would lead to large  
17 errors in the prediction of CCN. As discussed in the previous section, the  $D_{50}$  was more sensitive to the  
18  $\sigma_{s/a}$  at lower supersaturations, leading to a large deviation of the  $N_{CCN}$  from the measured value. The  
19 best agreement between the calculated AR and the measured AR was seen using scheme 5 as the slopes  
20 at the four SS were close to 1 (Fig. 10q-t).

21

### 1 3.3.2 The $N_{CCN}$ prediction based on AMS measurements

2 We proposed five approaches based on H-TDMA measurements to predict the  $N_{CCN}$  in the previous  
3 section. Alternatively, we can calculate the  $N_{CCN}$  based on AMS measurements. Here we proposed four  
4 methods based on either size-resolved chemical composition or bulk  $PM_{10}$  chemical composition from  
5 the AMS measurements (Table 5). Here we assumed that the particles were internally mixed and the  
6 median  $\kappa_{AMS}$  obtained from bulk composition was 0.28, higher than those from size-resolved  
7 composition (0.24-0.26 in Fig. 3), probably due to a higher mass fraction of inorganic matters in bulk  
8 NR- $PM_{10}$  (Fig. 2). We excluded the size-resolved data at 0.7% SS due to their poor quality. Note that the  
9 impact on the calculated  $\kappa_{AMS}$  values and the predicted  $N_{CCN}$  was minor using the  $\kappa$  value (0.53) of  
10 ammonium sulfate from Petters and Kreidenweis (2007). For example, the  $\kappa_{AMS}$  values slightly  
11 increased from 0.27 to 0.28 at 0.1% SS; the slopes for scheme 6, 8 and 9 in Table 5 slightly increased  
12 from 0.9859 to 0.9898, 0.9721 to 0.9834, and 0.9742 to 0.9973, respectively, while the one for scheme  
13 7 did not change. Figure 11 shows the correlation between the measured and predicted  $N_{CCN}$  from  
14 schemes 6-9. The  $N_{CCN}$  was under-predicted at 0.1% SS and was over-predicted at 0.7% SS. We  
15 proposed three potential factors that might impact  $N_{CCN}$  prediction based on AMS measurements. (1)  
16 The assumed  $\kappa_{org}$  values were probably underestimated for particles larger than 100 nm, leading to the  
17 underestimated  $N_{CCN}$  at low SS. As shown in Fig. 3, the predicted  $\kappa$  shows a larger deviation from the  
18 measured value for a larger particle. The  $D_{50}$  values were more sensitive to particle hygroscopicity at  
19 lower SS as discussed in the previous section. (2) The pure water assumption for surface tension. As we  
20 have shown in the previous section, the  $\sigma_{s/a}$  values for the aerosol particles were found to be much  
21 smaller than the  $\sigma_{s/a}$  for pure water ( $0.072 \text{ N m}^{-1}$ ). As a result, the pure water assumption for surface

1 tension led to the  $N_{CCN}$  underestimation. In addition, again the  $D_{50}$  was more sensitive to  $\sigma_{s/a}$  at the low  
2 SS. (3) The exclusion of black carbon (BC) particles and the mixing state assumption. The BC particles  
3 were known to be non-hygroscopic and had a low CCN activity. During the campaign period, the  
4 average BC concentration was about  $5.91 \mu\text{g}/\text{m}^3$  which accounts for 7 % in  $\text{PM}_{2.5}$ . The assumption of  
5 no BC particles would lead to the overestimation of  $N_{CCN}$ . Here, the particles were assumed to be  
6 internally mixed in the AMS measurements. This would lead to an overestimation of the  $N_{CCN}$  when the  
7 ambient particles tend to be externally mixed (Wang et al., 2010; Sánchez Gacita et al., 2017). However,  
8 the internal mixing assumption seems to play a minor role in predicting the  $N_{CCN}$  at 0.1% SS since the  
9 particles at about 140-180 nm tend to be internally mixed as shown in Fig. 5. In this case, the  $\kappa_{org}$   
10 assumption and the pure water assumption played more important roles than the mixing state  
11 assumption at low SS (i.e., 0.1% SS). Figure 11 shows significant underestimation of  $N_{CCN}$  at 0.1% SS  
12 (panels a, e, i, m), while more or less comparable to the measured  $N_{CCN}$  at higher SS (i.e. 0.2%, 0.4%,  
13 0.7%). The difference between the  $\kappa_{AMS}$  and  $\kappa_{CCN}$  became smaller and the corresponding  $D_{50}$  value  
14 decreased with the increase of the SS so that the impacts of the  $\kappa_{org}$  assumption and the pure water  
15 assumption became minor with the increase of the SS. Instead, the internal mixing state assumption  
16 would play a more important role in the prediction (Meng et al. 2014). As shown in Fig. 5, the peak  
17 height and area of the less-hygroscopic mode became larger for the smaller size particles (i.e, 40 nm  
18 particles), implying that small particles were likely to be externally mixed, that is, the non or less  
19 hygroscopic species including BC and insoluble organics were less likely coated with inorganics salts.  
20 Hence the internal mixing assumption could lead to an overestimated  $N_{CCN}$ .

21

1 As discussed above, the two important parameters ( $\kappa_{\text{org}}$  and  $\sigma_{\text{s/a}}$ ) had significant impacts on the  $N_{\text{CCN}}$   
2 prediction. We denoted  $\kappa_{\text{org}}^*$  and  $\sigma_{\text{s/a}}^*$  as important representations respectively for hygroscopicity and  
3 surface tension contributed from organics. We also pointed out that the  $\kappa_{\text{org}}$  was dependent on the  
4 particle size and hence here we further assumed the  $\kappa_{\text{org}}^*$  values to be 0.15 and 0.1 respectively for  
5 particles larger and smaller than 100 nm. Note that we previously assumed the  $\kappa_{\text{org}}$  to be 0.1 for all  
6 particle sizes. Here we gave an example of the improvements at 0.1% SS when the  $\kappa_{\text{org}}$  and  $\sigma_{\text{s/a}}$  values  
7 were respectively replaced with the  $\kappa_{\text{org}}^*$  and  $\sigma_{\text{s/a}}^*$  ones (Fig. 12). The  $\kappa_{\text{AMS}}$  value calculated at 0.1% SS  
8 based on  $\kappa_{\text{org}}^*$  was 0.288, very close to the corresponding  $\kappa_{\text{CCN}}$  value (0.30), indicating that an  
9 improvement was made for the  $N_{\text{CCN}}$  prediction when including the  $\kappa_{\text{org}}^*$  value. The  $N_{\text{CCN}}$  prediction  
10 could be greatly improved when include both  $\sigma_{\text{s/a}}^*$  and  $\kappa_{\text{org}}^*$  in the calculation. For example, the  
11 underestimate of  $N_{\text{CCN}}$  decrease from 44% (Fig. 11a) to 4% (Fig. 12b). In addition, we also investigated  
12 the effects of the  $\sigma_{\text{s/a}}$  values in a range of 0.054 to 0.062  $\text{N m}^{-1}$  as discussed in section 3.2. The shadow  
13 area in Fig. 12b represents the variation of linear fit between the measured and predicted  $N_{\text{CCN}}$ . An  
14 under- and over-estimated value of 16% ( slope=0.84) and 8% (slope=1.08) was obtained for the  
15 predicted  $N_{\text{CCN}}$  to the measured  $N_{\text{CCN}}$  using a  $\sigma_{\text{s/a}}$  value of 0.054 and 0.062  $\text{N m}^{-1}$  respectively,  
16 indicating that the predicted  $N_{\text{CCN}}$  agreed reasonably with the measured ones when the  $\sigma_{\text{s/a}}$  values  
17 between 0.054 and 0.062  $\text{N m}^{-1}$  were used in this study. We conclude that the predicted  $N_{\text{CCN}}$  can agree  
18 better with the measured one when including both  $\sigma_{\text{s/a}}^*$  and  $\kappa_{\text{org}}^*$  in the calculation at low SS.

19

#### 20 **4 Summary and Conclusions**

21 The CCN activity is an important parameter that determines the extent to which atmospheric particles

1 can influence cloud formation. It is hence essential to predict CCN activity so that a quantitative  
2 assessment of atmospheric particles on cloud formation can be made. While numerous studies were  
3 performed to investigate the CCN activity under different atmospheric conditions around the world,  
4 only a few of them were made in the PRD region in China. In this study, several advanced instruments  
5 (i.e., the SMCA, AMS and H-TDMA) were used to respectively measure CCN activity, chemical  
6 composition and hygroscopicity in PRD during wintertime 2014. A variety of schemes were proposed  
7 to determine the CCN activity based on the measurements. Here two important properties were  
8 considered when evaluating the CCN activity: the hygroscopic parameter  $\kappa$  and the surface tension of  
9 the particles. Three methods (i.e., the SMCA, the AMS+ZSR, and the H-TDMA) were employed to  
10 calculate the  $\kappa$  values based on our measurements. The results show that the deviation between  $\kappa_{\text{AMS}}$   
11 and  $\kappa_{\text{CCN}}$  became larger at low supersaturation ratios, indicating that the organic components in larger  
12 size particles were more aged and hygroscopic. The activation curve became smoother at the low SS,  
13 which could be partly attributed to the higher heterogeneity of chemical composition for larger particles.  
14 In general, the Gf-PDF measured by H-TDMA exhibited a bimodal distribution with a less-hygroscopic  
15 mode and a more-hygroscopic mode. The less-hygroscopic mode was more significant at smaller  
16 diameters, indicating a more external mixing for smaller particles, while the more-hygroscopic mode  
17 increased with diameter and became broader, implying higher hygroscopicity and more complex  
18 chemical composition for larger particles. The shapes of activation curve were related to the  $\sigma$  values of  
19 the Gf-PDF. The higher  $\sigma$  values suggest the higher heterogeneity of chemical composition and smooth  
20 activation curve. A  $\kappa$  value of 0.22-0.30 measured by H-TDMA was obtained for 40-200 nm particles in  
21 this study during the measurement period, larger than those previously measured in the PRD region,

1 indicating an increasing mass fraction of nitrate in recent years.

2

3 Organic compounds could influence CCN activity through modifying the hygroscopicity and surface  
4 tension of the particles. The impacts of organics on CCN activity were also investigated in this study.  
5 The increase of organic mass fraction in the particles could lead to the decrease of the aerosol  
6 hygroscopicity and hence increase the  $D_{50}$ , especially at low supersaturation. In addition, organics  
7 could decrease the surface tension  $\sigma_{s/a}$ . This could lead to the underestimated CCN activity if pure  
8 water solution is assumed when inverting the H-TDMA data. We evaluated the impact of the surface  
9 tension on the activation ratios over a wide range of  $\sigma_{s/a}$  values (0.03-0.07 N m<sup>-1</sup>) for several measured  
10 size particles (40, 80, 110, and 150 nm) and found that a  $\sigma_{s/a}$  value of 0.058 N m<sup>-1</sup> was the best fit  
11 between predicted AR and measured AR, which could then be used to predict the CCN activity in the  
12 PRD region. Based on the hygroscopicity and chemical composition measured in this study, we  
13 proposed several schemes to predict the CCN activity. Overall, the predicted  $N_{CCN}$  agreed well with the  
14 measure one. The slope and  $R^2$  of  $N_{CCN}$  predicted from average data was similar to the  $N_{CCN}$  predicted  
15 from real time data. The  $N_{CCN}$  obtained from H-TDMA measurement was under-predicted, if pure  
16 water assumption was used and better agreement with the measured values can be achieved by using  
17 the adjusted  $\sigma_{s/a}$  (i.e.,  $\sigma_{s/a}^* = 0.058$  N m<sup>-1</sup>). Similarly, the  $N_{CCN}$  predicted from AMS measurement was  
18 underestimated at low supersaturations and overestimated at high supersaturations, due to an  
19 assumption of fixed 0.1 for  $\kappa_{org}$  and the external mixing state. Better predicted CCN concentrations can  
20 be obtained by using  $\sigma_{s/a}^*$  and  $\kappa_{org}^*$  in the calculation, especially at low supersaturations. For high  
21 supersaturations, the effect of internal mixing assumption should be taken into consideration. We

1 concluded that better agreement between predicted and measured CCN concentrations could be  
2 achieved by taking the effects of organic into account on the hygroscopicity, surface tension, and the  
3 mixing state of the particles. More work on the roles of organics on the CCN activity is obviously  
4 needed in order to better understand the impacts of atmospheric particles on cloud formation and hence  
5 climate.

6

### 7 **Acknowledgement**

8 The authors acknowledge the support from the following funding agencies: National Key R&D  
9 Program of China (2016YFC0201901, 2017YFC0209500, 2016YFC2003305), National Natural  
10 Science Foundation of China (NSFC) (91644225, 21577177, 41775117), and Guangdong provincial  
11 scientific planning project (2014A020216008, 2016B050502005). J.Z. also acknowledges the funding  
12 from the “111 Plan” Project of China (Grant B17049).

13

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11

1 **Table 1.** The  $\kappa$  values of the related species in the study.

Species	$\kappa$
$\text{NH}_4\text{NO}_3$	0.58 <sup>a</sup>
$\text{NH}_4\text{HSO}_4$	0.56 <sup>a</sup>
$\text{H}_2\text{SO}_4$	0.90 <sup>a</sup>
$(\text{NH}_4)_2\text{SO}_4$	0.48 <sup>a</sup>
Organics	0.10 <sup>b</sup>

2 <sup>a</sup>  $\kappa$  of inorganics compounds are derived from ADDEM (Topping et al., 2005)

3 <sup>b</sup>  $\kappa$  of organics is from Meng et al. (2014)

4

5

1 **Table 2.** Summary of the measured CCN concentration, activation ratio, and  $D_{50}$  at the four  
 2 supersaturations during the campaign.

SS		0.1%	0.2%	0.4%	0.7%
	Max	15165	19989	25964	26208
$N_{CCN}(\#/cm^3)$	Min	258	361	408	502
	Mean $\pm$ STD	3103 $\pm$ 1913	5095 $\pm$ 2972	6524 $\pm$ 3783	7913 $\pm$ 4234
	Max	0.68	0.75	0.89	0.94
$N_{CCN}/N_{CN,tot}$	Min	0.06	0.10	0.19	0.28
	Mean $\pm$ STD	0.26 $\pm$ 0.10	0.41 $\pm$ 0.14	0.53 $\pm$ 0.15	0.64 $\pm$ 0.13
	Max	268.90	194.04	145.28	97.17
$D_{50}$ (nm)	Min	112.47	76.60	43.50	24.21
	Mean $\pm$ STD	156.02 $\pm$ 19.48	106.66 $\pm$ 16.99	77.96 $\pm$ 14.86	58.45 $\pm$ 10.68

3

1 **Table 3.** The average of  $\sigma$  values of Gf-PDF measured by H-TDMA for five diameters for the steep and  
 2 smooth activation ratio at four supersaturations.

SS(%)	0.1		0.2		0.4		0.7	
Dp(nm)	Steep	Smooth	Steep	Smooth	Steep	Smooth	Steep	Smooth
40	0.13	0.17	0.12	0.17	0.11	0.19	0.11	0.19
80	0.16	0.20	0.14	0.20	0.14	0.21	0.14	0.20
110	0.17	0.21	0.15	0.21	0.15	0.21	0.16	0.20
150	0.19	0.22	0.17	0.23	0.17	0.22	0.18	0.21
200	0.20	0.23	0.19	0.24	0.19	0.24	0.19	0.23

3

1 **Table 4.** The overview of different schemes used in the  $N_{CCN}$  prediction based on H-TDMA  
2 measurements.

<b>Scheme</b>	<b>Method</b>	<b>Slope</b>	<b>R<sup>2</sup></b>
1	Real time activation curve	0.8275	0.93
2	Average activation curve	0.8183	0.93
3	Real time $D_{50}$	0.8869	0.93
4	Average $D_{50}$	0.8738	0.93
5	Real time activation curve using $\sigma_{s/a}^*$	0.9377	0.93

3

1 **Table 5.** The overview of methods used in the  $N_{CCN}$  prediction based on AMS measurements.

<b>Scheme</b>	<b>Method</b>	<b>Slope</b>	<b>R<sup>2</sup></b>
6	Real time bulk composition	0.9859	0.91
7	Average bulk composition	1.0108	0.91
8	Real time size-resolved composition	0.9721	0.87
9	Average size-resolved composition	0.9742	0.86

2

1 FIGURE CAPTIONS

2 Fig. 1. A schematic representation of  $N_{CCN}$  prediction based on the H-TDMA and the AMS  
3 measurements. The  $N_{CCN}$  can be predicted based on the fitted activation ratio (approach I) and the  $D_{50}$   
4 (approach II) both obtained from the H-TDMA measurement, the size-resolved composition (approach  
5 III) and the bulk  $PM_{10}$  composition (approach IV) both obtained from the AMS measurements. Panel (a)  
6 is the representation of calculating the activation ratio for a specific diameter and SS and the shadow  
7 area represents the particles which can be activated as CCN; (b) and (c) are the representations of the  $\kappa$   
8 values obtained respectively from size-resolved chemical composition and bulk chemical composition;  
9 (d) is the representation of fitting the activation ratio to the particle diameter  $D_p$  (red dot); (e), (f), (g),  
10 and (h) are the representations of predicting the  $N_{CCN}$  using the four approaches respectively and the  
11 shadow area represents the particles which can be activated as CCN.

12 Fig. 2. The mass fraction of the bulk NR- $PM_{10}$  composition (a) and the mass fraction of the  
13 size-resolved composition (b).

14 Fig. 3. The median and interquartile PNSD,  $\kappa$  obtained from H-TDMA, AMS and CCN measurement  
15 during the campaign. The  $\kappa$  was plotted against their corresponding median  $D_{50}$  (SMCA and AMS) or  
16 measured diameter (H-TDMA). Dot points represent the median value and the bars represent the  
17 interquartile range. The blue, red, and green represent  $\kappa_{CCN}$ ,  $\kappa_{AMS}$  and  $\kappa_{HTDMA}$  respectively.

18 Fig. 4. The sized resolved activation ratios measured by the SMCA at four different supersaturations.  
19 Note that the curves were fitted according to the SMCA measurements.

20 Fig. 5. The Gf-PDF as a function of Gf measured by H-TDMA for the five diameters

21 Fig. 6. The relationship between size-resolved mass fraction of organics and  $D_{50}$  at three



1 supersaturations. The red, blue, and green dots and line represent 0.1%, 0.2%, and 0.4% SS  
2 respectively.

3 Fig. 7. The predicted activation ratio based on H-TDMA measurement vs. the measured activation ratio  
4 at 0.1%, 0.2%, 0.4% and 0.7% SS for 40, 80, 110, 150 and 200 nm particles.

5 Fig. 8. The relative deviation between predicted AR and measured AR at different assumed  $\sigma_{s/a}$ ; the  
6 color code represents  $R^2$  between calculated AR and measured AR.

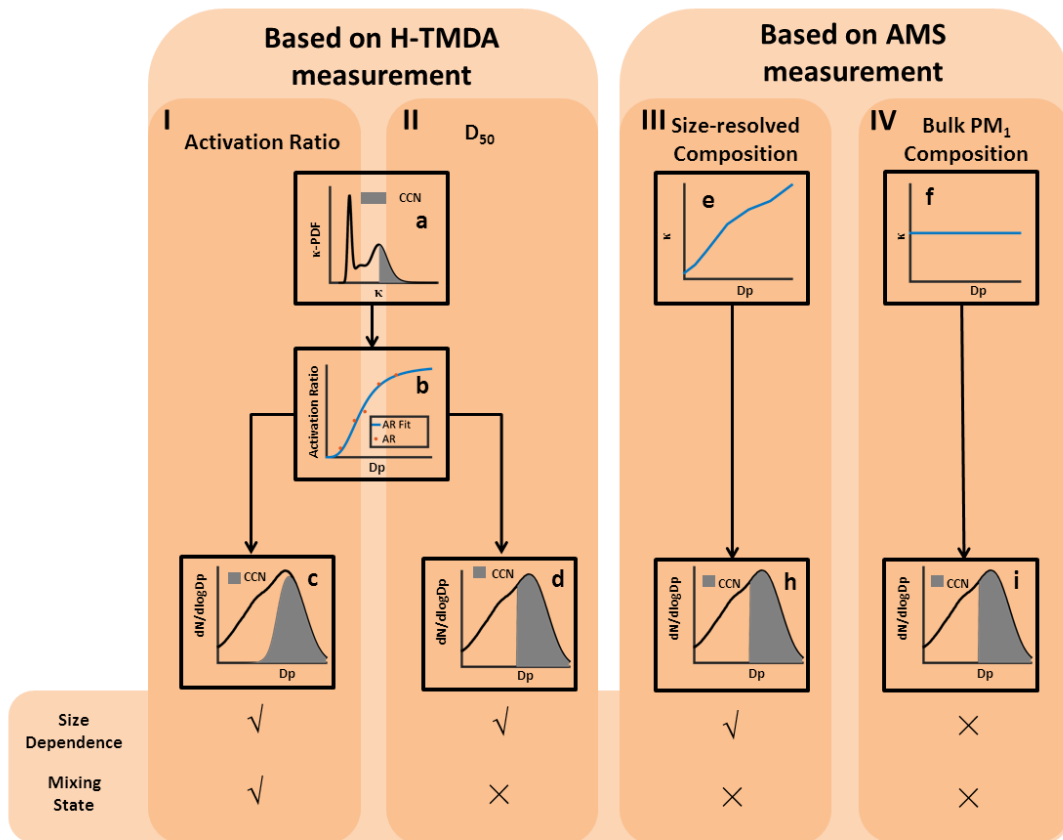
7 Fig. 9. The predicted activation ratio using new surface tension assumption ( $\sigma_{s/a}^*$ ) based on H-TDMA  
8 measurement vs. the measured activation ratio at 0.1%, 0.2%, 0.4% and 0.7% SS for 40, 80, 110, 150  
9 and 200 nm particles.

10 Fig. 10. The relationship between measured  $N_{CCN}$  and predicted  $N_{CCN}$  based on scheme 1, 2, 3, 4 and 5.  
11 The black lines represent 1:1 lines.

12 Fig. 11. The relationship between measured  $N_{CCN}$  and predicted  $N_{CCN}$  based on scheme 6, 7, 8 and 9.  
13 The black lines represent 1:1 lines.

14 Fig. 12. The relationship between measured  $N_{CCN}$  and predicted  $N_{CCN}$  at SS 0.1% based on  
15 size-resolved chemical composition using  $\kappa_{org}^*$  (a) and  $\kappa_{org}^*$  and  $\sigma_{s/a}^*$  (b). The shadow area represents  
16 the variation of the linear fit using the  $\sigma_{s/a}$  values from 0.054 to 0.062 N m<sup>-1</sup>.

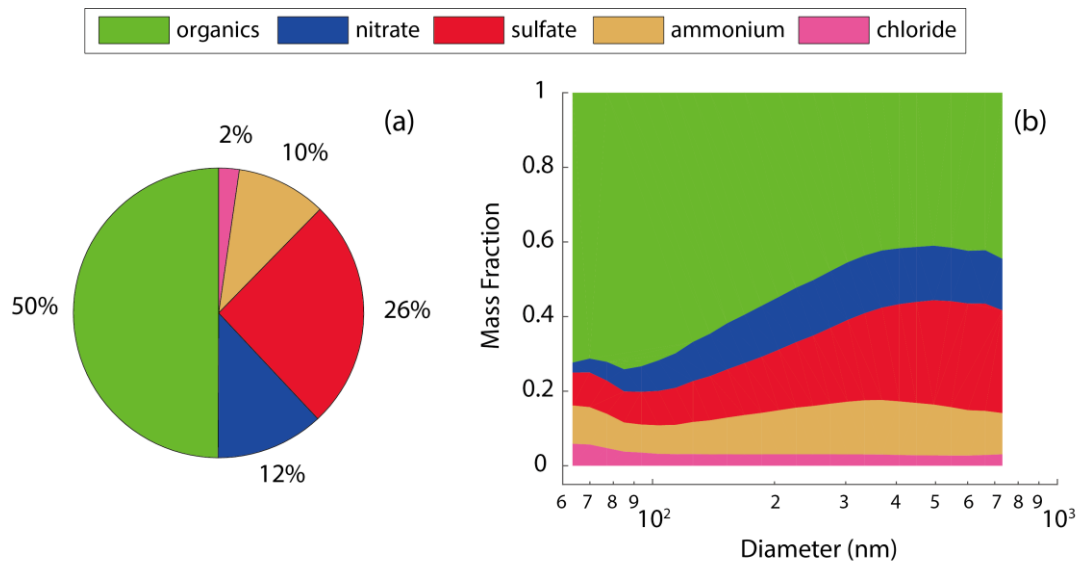
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Fig. 1.

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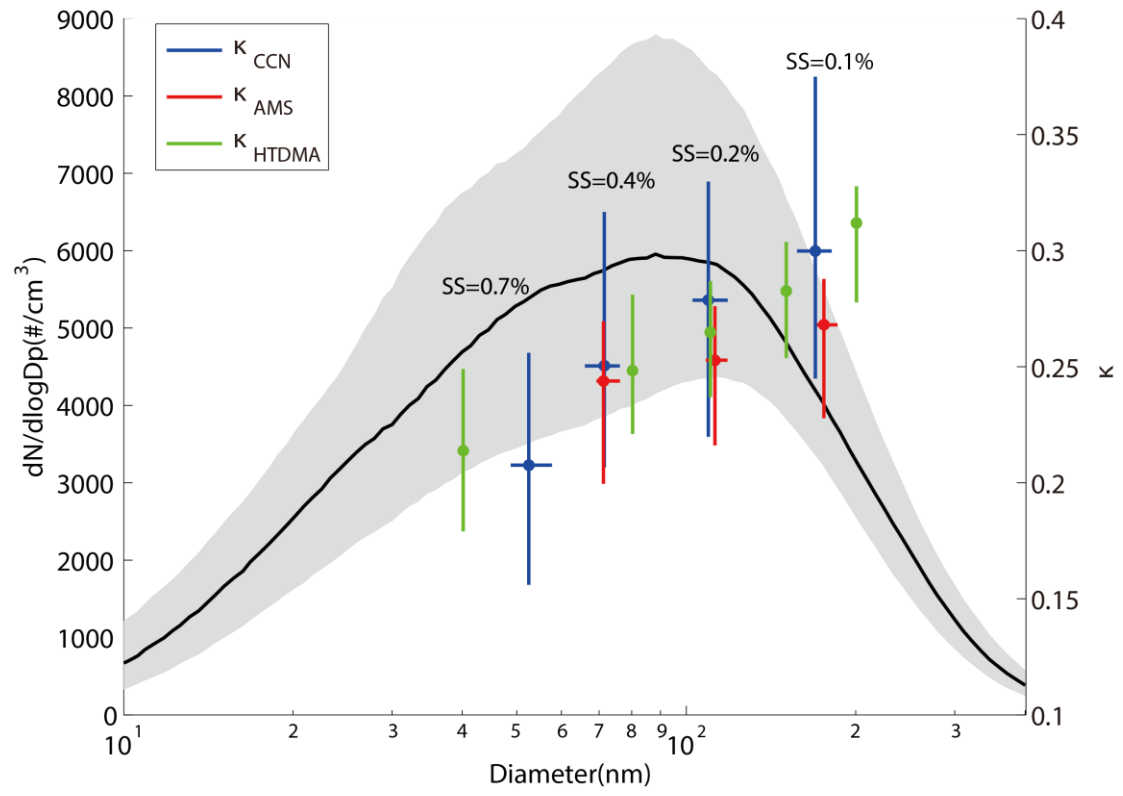
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Fig. 2.

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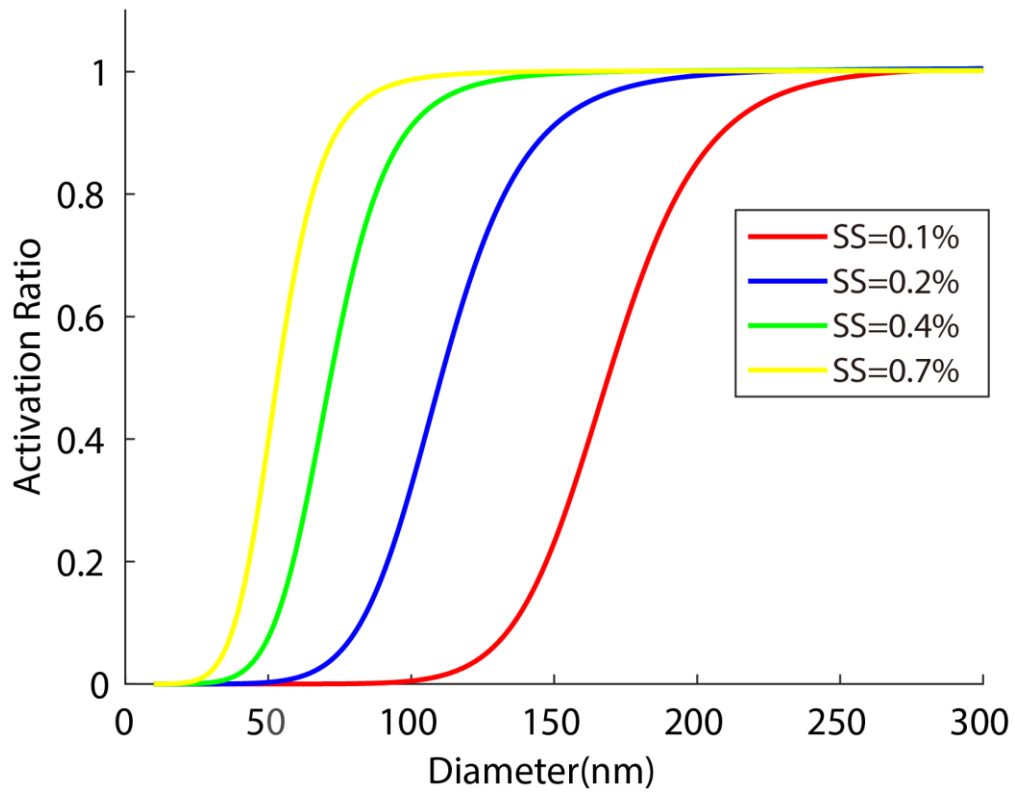
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2 Fig. 3.

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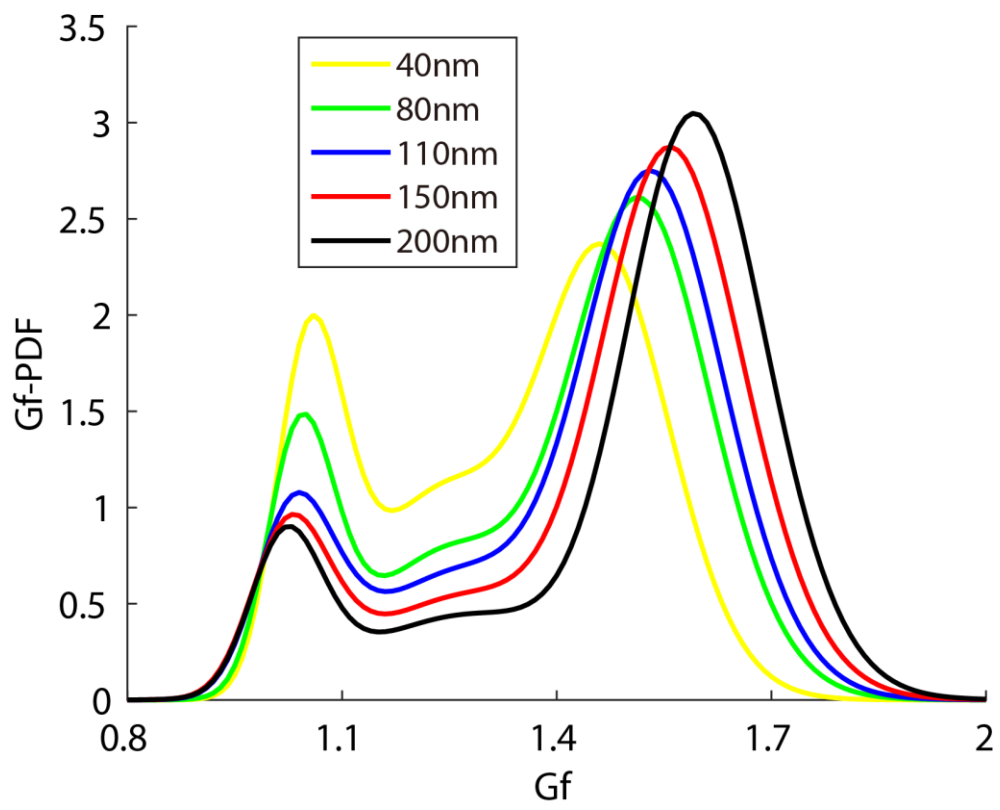
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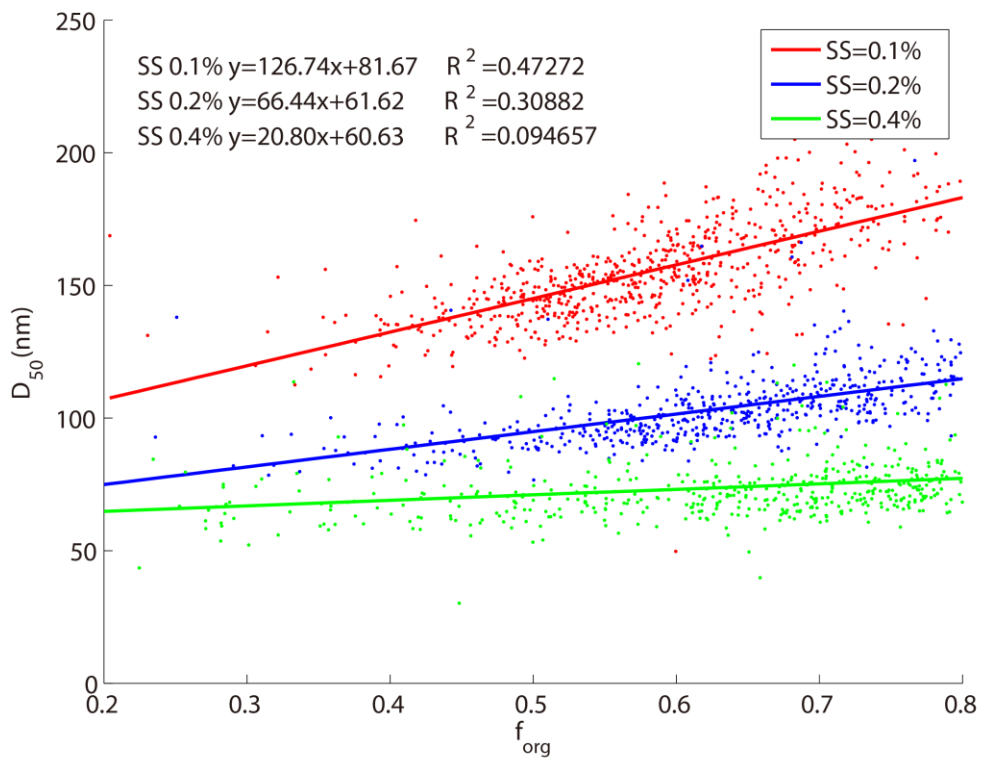
3 Fig. 4.

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Fig. 5.

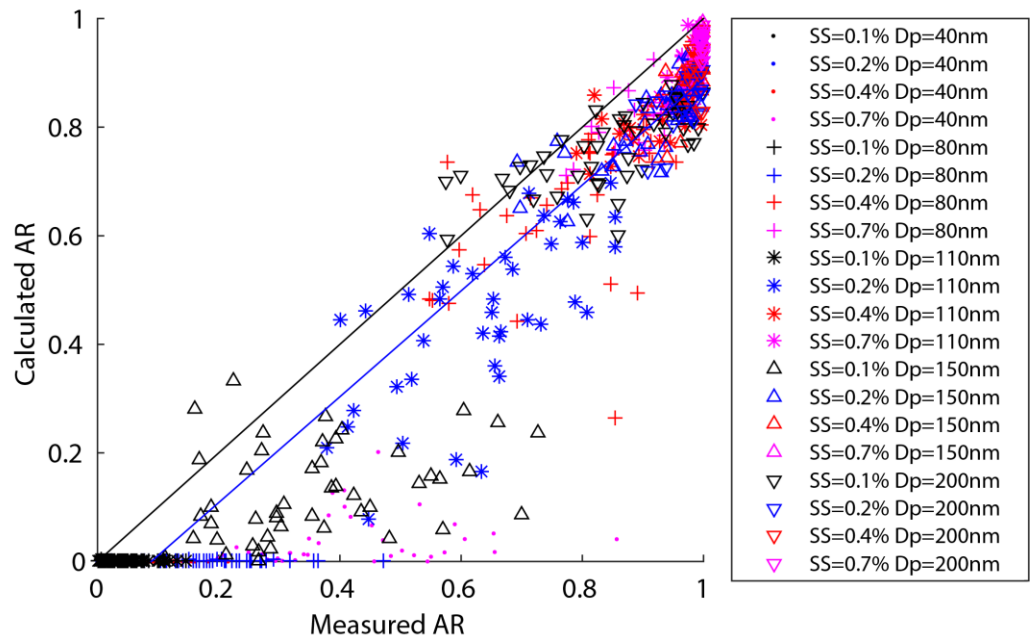


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2 Fig. 6.

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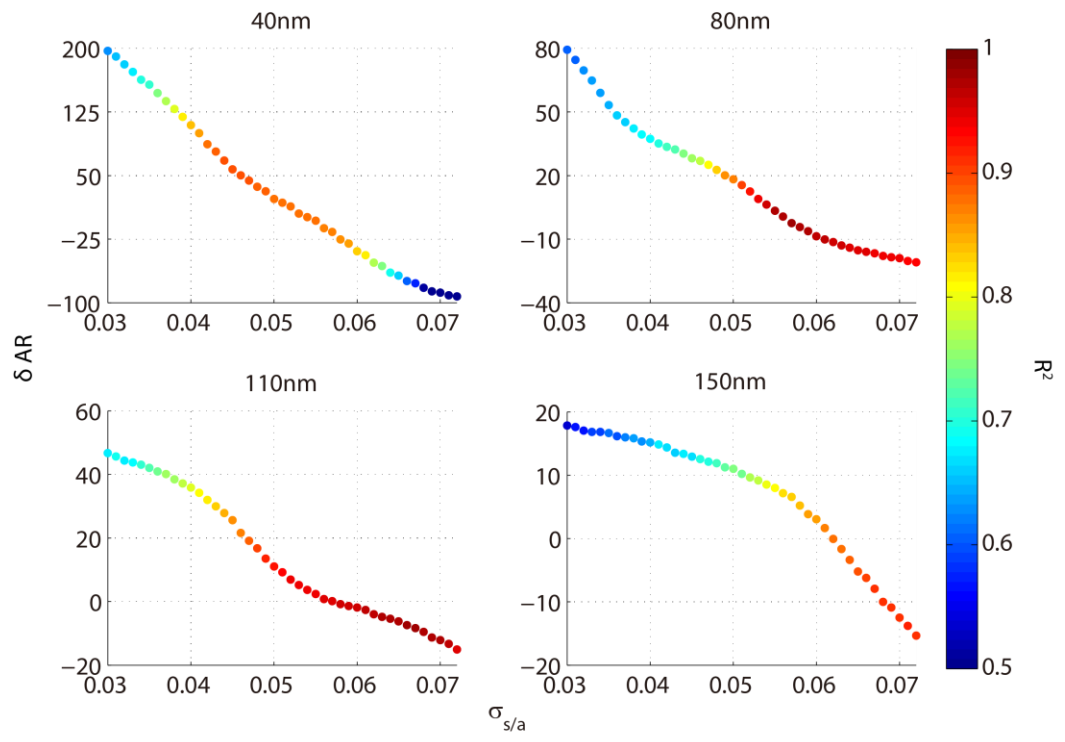
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3 Fig. 7.

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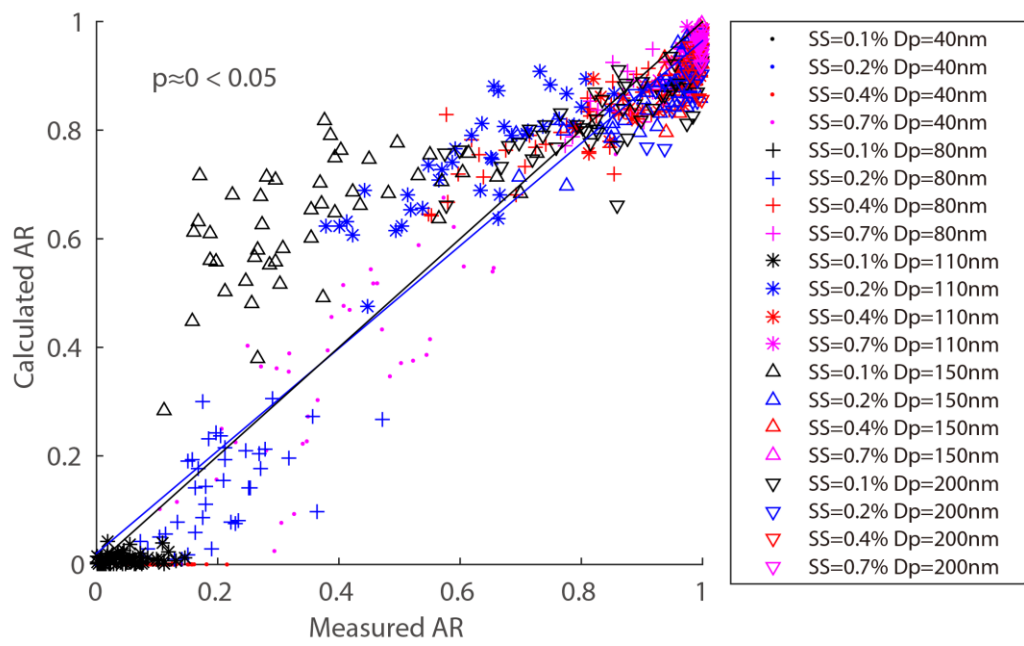
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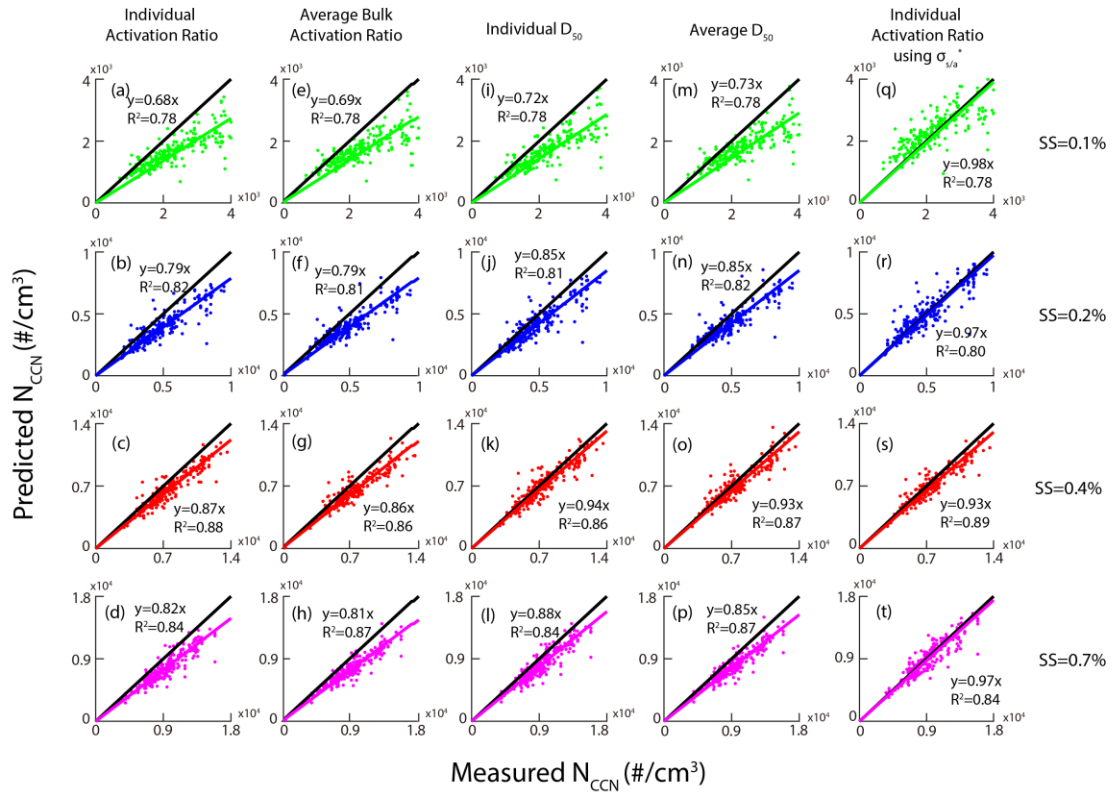
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2 Fig. 9.

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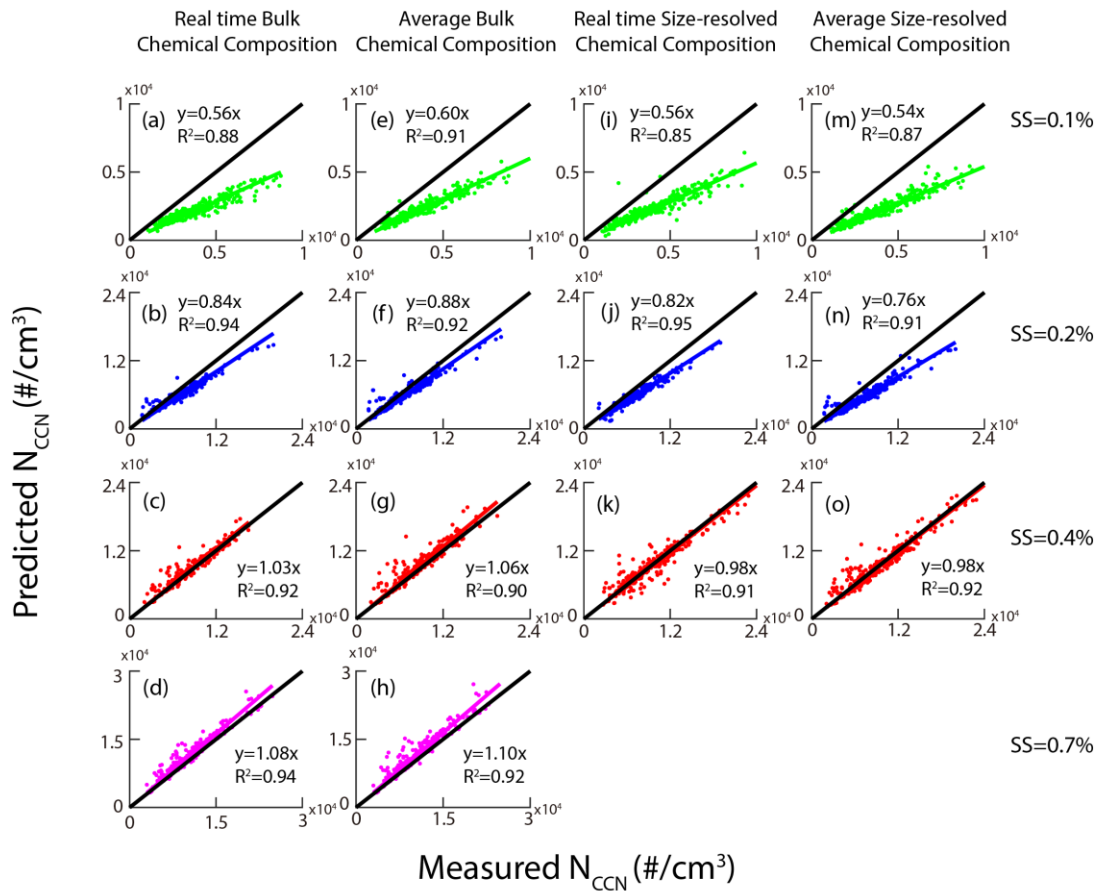


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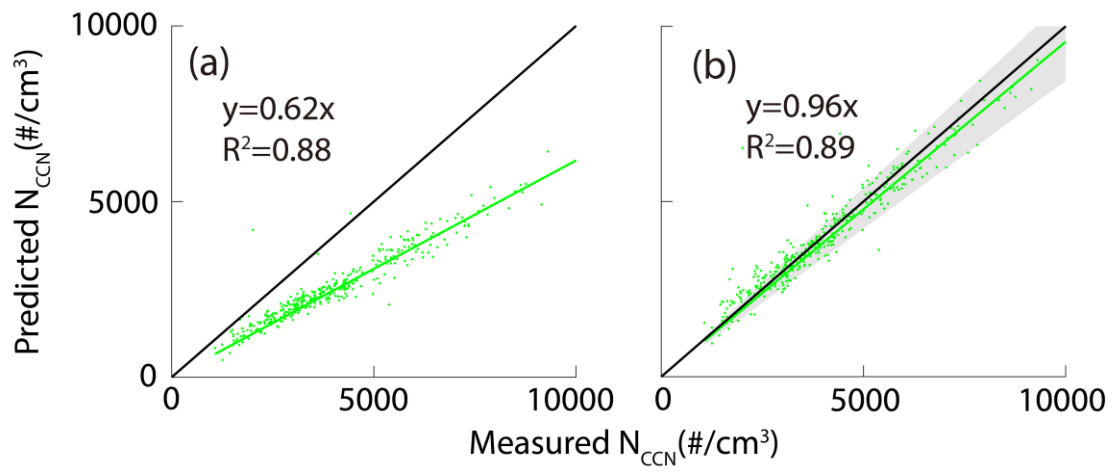
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2 Fig. 11.

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3 Fig. 12.