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2	The Influence of HCl on the Evaporation Rates of $H_2O$ over Water Ice in
3	the Range 188 to 210 K at small Average Concentrations
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19	ABSTRACT
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20 21 22	The evaporation flux $J_{ev}(H_2O)$ of $H_2O$ from HCl-doped typically 1.5 µm or so thick vapor- deposited ice films has been measured in a combined quartz crystal microbalance (OCMB) – residual
23	gas mass spectrometry (MS) experiment. $J_{ev}(H_2O)$ has been found to show complex behaviour and to
24	be a function of the average mole fraction $\chi_{HCl}$ of HCl in the ice film ranging from $6 \times 10^{14}$ to $3 \times 10^{17}$
25 26	molecule cm <sup>2</sup> s <sup>-</sup> at $1/4 - 210$ K for initial values $\chi^{2}_{HCl}$ ranging from $5x10^{\circ}$ to $3x10^{\circ}$ at the start of the evaporation. The dose of HCl on ice was in the range of 1 to 40 formal monolayers and the H <sub>2</sub> O vapor
27	pressure was independent of $\gamma_{HCI}$ within the measured range and equal to that of pure ice down to 80
28	nm thickness. The temporal dependence of $J_{ev}(H_2O)$ was correlated with (a) the evaporation range $r^{b/e}$
29	as the ratio of $J_{ev}(H_2O)$ just before HCl-doping of the pure ice film and $J_{ev}(H_2O)$ after observable HCl
30	desorption towards the end of film evaporation, and (b) the remaining thickness $d_{\text{D}}$ below which
31	$J_{ev}(H_2O)$ decreases to less than 85% of pure ice. The time dependence of $J_{ev}(H_2O)$ from HCl-doped ice
52 22	This suggests two limiting data sets, one associated with the occurrence of a two-phase pure
33 34	HCl/H <sub>2</sub> O binary mixture (set B). The measured values of L <sub>2</sub> (H <sub>2</sub> O) may lead to significant evaporative
35	life-time extensions of HCl-contaminated ice cloud particles under atmospheric conditions, regardless

36 of whether the structure corresponds to an amorphous or crystalline state of the HCl/H<sub>2</sub>O aggregate.





### 38 1. INTRODUCTION

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40 HCl is among the mineral acids that control the acidity of the atmosphere together 41 with  $HNO_3$  and  $H_2SO_4$ . The production of atmospheric HCl is predominantly taking place in 42 the middle and upper stratosphere where  $O_3$  is produced owing to photolysis of halogen 43 containing source gases such as CFC's (chlorofluorocarbons). However, there are no known 44 sources of HCl in the upper troposphere (UT) because scavenging processes of HCl 45 throughout the troposphere are very efficient which leads to HCl background concentrations of less than 0.1 ppb (Graedel and Keene, 1995). The absence of significant sources in the 46 47 troposphere, the long photolytic lifetime of HCl and the fact that the production region is well 48 separated from the regions of interest, namely the UT and the lower stratosphere (LS) all 49 contribute to the fact that HCl is an excellent tracer for stratospheric ozone in the UT (Marcy 50 et al., 2004). Owing to the frequent occurrence of Cirrus clouds in this atmospheric region it 51 is of obvious interest to study the interaction of HCl with atmospheric ice particles at relevant 52 temperature and pressure conditions (Jensen et al. 2001; Zerefos et al., 2003). The compact 53 correlation between O<sub>3</sub> and HCl has been used to monitor stratospheric-tropospheric 54 exchange processes and stratospheric O<sub>3</sub> intrusions into the troposphere that are still an active 55 field of investigation (Houghton et al, 2001).

HCl is of importance in the LS as it partakes in heterogeneous reactions on Polar
Stratospheric Ice Clouds (PSC's) as well as on background stratospheric H<sub>2</sub>SO<sub>4</sub> aerosol
according to the following reaction taken as an example:

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## $HCl(ads) + ClONO_2 \rightarrow Cl_2(g) + HNO_3(ads)$

60 These reactions efficiently convert inactive Cl-containing reservoir molecules such as HCl 61 and  $CIONO_2$  into active photolyzable Cl-containing compounds in a single reaction. Typical examples of such photolabile reaction products are Cl<sub>2</sub>, CINO<sub>2</sub> and HOCl that will change the 62 atmospheric composition owing to the high reactivity of the photolysis products such as 63 64 atomic Cl (Solomon et al., 1986; Tolbert et al., 1987; WMO 2003). It thus follows that HCl is 65 of stratospheric importance and is frequently used as a model compound for heterogeneous 66 reactions on ices that has inspired many laboratory kinetic studies (Leu et al., 1991; Hanson and Ravishankara, 1992; Chu et al., 1993; Flückiger et al., 1998; Hynes et al., 2001; Abbatt, 67 2003). 68

HCl forms hydrates of variable stoichiometry when exposed to ice depending on the
temperature of deposition and the partial pressure of HCl (Ortega et al., 2004; Graham and
Roberts, 1997). X-Ray diffraction has allowed the identification of four crystalline hydrates





72 containing one (Yoon and Carpenter, 1959), two (Lundgren and Olovson, 1967), three 73 (Lundgren and Olovson, 1967a) and six (Taesler and Lundgren, 1978) H<sub>2</sub>O per HCl molecule. 74 In addition, amorphous mono-, tetra- and hexahydrates have been reported under various 75 experimental conditions (Yoon and Carpenter, 1959; Delzeit et al., 1993). The control of 76 growth conditions of a specific HCl hydrate is sometimes elusive, but the formation of a 77 saturated HCl hexahydrate phase has been reported at sufficiently large HCl exposure 78 (Graham and Roberts, 1995) using amorphous ice as a starting point despite the fact that the 79 hexahydrate is said to nucleate with difficulty, at least in thin films (Ortega et al., 2004). 80 However, the molecular and dynamic details of the crystallization process have not been 81 investigated as yet.

82 Fourier-Transform IR (FTIR) absorption measurements have enabled the 83 characterization of both amorphous as well as crystalline HCl hydrates at growth conditions 84 that are sometimes significantly different compared to the samples investigated using X-Ray 85 diffraction. Vibrational spectra of HCl hydrates in the mid IR have been routinely used for identification purposes for some time (Ferriso and Hornig, 1955; Gilbert and Sheppard, 86 87 1973). Recently, the mid IR absorption spectra of the four HCl hydrates mentioned above 88 have been assigned in a comprehensive and definitive way, albeit without simultaneous proof 89 of the crystalline structure using X-Ray diffraction (Buch et al., 2002; Xueref and Dominé, 90 2003). More recently, the reflection absorption IR spectrum (RAIR) of crystalline HCl 91 hexhydrate in the mid-IR range has been recorded and assigned using theoretical calculations 92 based on density functional theory that results in a refinement of the geometric structure of 93 the HCl hydrates and a prediction of the vibrational modes of the crystal (Ortega et al., 2004).

94 Regarding the nature of the HCl-ice adsorbate one of the important questions is 95 whether adsorbed HCl is ionized or exists as a molecular adsorbate under atmospherically relevant conditions of the UT/LS. This will determine the mechanism of the heterogeneous 96 97 reaction which constitutes necessary knowledge for the extrapolation of heterogeneous 98 reaction rates measured in the laboratory to atmospheric conditions. Thermal desorption of 99 HCl monitored by IR absorption in the mid-IR range revealed a molecularly adsorbed state of 100 HCl desorbing below 50 K (Delzeit et al., 1993a). IR studies performed by Banham et al. on 101 HCl-ice films failed to detect molecularly adsorbed HCl at  $T \ge 90$  K despite the high rate of 102 HCl adsorption in that temperature range (Banham et al., 1995). In contrast, Graham and 103 Roberts attributed a characteristic Temperature Programmed Desorption (TPD) peak of a 104 HCl/amorphous ice adsorbate monitored by residual gas MS and occurring at 150 K to 105 molecularly adsorbed HCl (Graham and Roberts, 1995). However, they did not report the IR





106 absorption spectrum of the adsorbate in the mid-IR nor did they explain why molecular 107 adsorption of HCl exclusively occurred on amorphous, but not on crystalline ice. Most recent 108 results seem to point towards the existence of molecularly adsorbed HCl on ice below 50 K 109 and at submonolayer coverages in coexistence with ionized solvated HCl whose fraction 110 increases with increasing ice temperature (Buch et al., 2002; Delzeit et al, 1993a; Delzeit et al. 111 1997; Uras et al., 1998; Devlin et al., 2002; Lu and Sanche, 2001). Kang et al. discovered 112 that both molecularly adsorbed as well as ionized HCl coexisted on ice that was deposited 113 under Ultra-High Vacuum (UHV) conditions in the temperature range 50 to 140 K and under 114 conditions of low HCl exposure (Kang et al., 2000).

115 Although theoretical electronic structure calculations predict spontaneous ionization of 116 adsorbed HCl (Gertner and Hynes, 1996; Bolton and Petterson, 2001) most experiments point 117 towards a seemingly thermally activated ionization process that may be enabled by structural 118 factors of the ice matrix that are themselves a function of temperature. Consistent with these 119 results concentration profiling experiments of HCl/ice adsorbates using static Secondary 120 Ionization Mass Spectroscopy (SIMS) techniques failed to discover molecularly adsorbed HCl on ice in the range 90-150 K (Donsig and Vickerman, 1997). In conclusion, both 121 122 experimental and theoretical studies clearly point to the absence of significant quantities of 123 molecularly or covalently adsorbed HCl under stratospheric conditions. Instead, HCl is 124 ionized and solvated by H<sub>2</sub>O on the surface of ice films and may occur either as amorphous 125 HCl/H<sub>2</sub>O hydrates of undefined stoichiometry or as crystalline HCl hydrates. However, these 126 facts do not rule out the presence of small amounts of molecularly adsorbed HCl on ice that 127 may be intermediates in the complex mechanism of HCl adsorption on ice as evidenced by 128 the negative temperature dependence of the rate of uptake of HCl on ice (Flückiger et al., 129 1998). In fact, such an intermediate has been invoked in the description of HCl adsorption on 130 ice under atmospheric conditions using a chemical kinetic model based on a multitude of 131 experimental observables collected upon HCl uptake on ice (Flückiger and Rossi, 2003).

132 Work by Parent and coworkers uses Near-Edge X-Ray Absorption Spectroscopy 133 (NEXAFS) of HCl-doped low temperature ice substrates in order to distinguish between bulk 134 and HCl surface states in the temperature range 20 to 150 K (Bournel et al., 2002; Parent and Laffon, 2005). NEXAFS is a sensitive method and nicely allows the distinction between 135 136 molecularly adsorbed HCl and the ionized form  $H_3O^+Cl^-$  on the one hand, and on the other hand it enables the distinction of HCl between the surface and the bulk of the ice substrate. 137 138 The results seem to confirm the consensus on the low-temperature existence of the 139 molecularly adsorbed HCl up to 90 K beyond which more and more HCl spontaneously





140 ionizes such that the process is complete at 150 K. More importantly, the results indicate that 141 the "dangling bonds" of the ice surface attributed to isolated OH groups are not the unique site of HCl adsorption, even in the low T range 20-90 K. This pioneering work suggests that 142 143 maiden uptake of HCl onto pure ice weakens and perturbs the crystal structure of the ice 144 matrix in an irreversible way such that additional sites for HCl adsorption and ionization are 145 created. Initial HCl uptake on pure ice therefore has a catalytic effect on the following HCl 146 uptake. This irreversible nature of initial HCl dosing has been observed some time ago in Knudsen flow reactor studies on the key heterogeneous reaction discussed above owing to the 147 148 persistent HCl partial pressure at temperatures representative of the UT/LS (Oppliger et al., 149 1997).

150 We have concluded from recent work that HCl doping in quantities of submonolayer 151 to several monolayers of HCl leads to the decrease of both the evaporative flux  $J_{ev}$  (molecule  $cm^{-2}s^{-1}$ ) or rate  $R_{ev}$  (molecule  $cm^{-3}s^{-1}$ ) and the rate of condensation  $k_{cond}$  (s<sup>-1</sup>), of H<sub>2</sub>O in the 152 presence of ice without perturbing the equilibrium vapour pressure of H<sub>2</sub>O, P<sub>H2O</sub><sup>eq</sup> (Delval et 153 154 al., 2003). We have furthermore shown that the way  $J_{ev}$  of  $H_2O$  decreases with time depends on the rate of deposition or the integral of deposited HCl, namely  $R_{HCl}$  (molecule s<sup>-1</sup>) and  $N_{HCl}$ 155 156 (molecule), respectively. It appears that two observed HCl species on/in ice, namely single 157 phase amorphous HCl/H<sub>2</sub>O mixtures and a binary phase consisting of pure ice and an as yet 158 unidentified crystalline HCl hydrate, HCl $\cdot$ xH<sub>2</sub>O, decrease  $J_{ev}(H_2O)$  to a different extent as 159 proposed in Delval et al. (2003). These results have led us to perform systematic experiments 160 using the Quartz Crystal MicroBalance (QCMB) combined with residual gas Mass 161 Spectrometry (MS) that we have used successfully in the past (Delval and Rossi, 2004) in 162 order to investigate the temporal change of  $J_{ev}(H_2O)$  with the increasing average mole fraction 163 of HCl,  $\chi_{HCl}$ , remaining in the ice. One of the goals of the present work is to determine the 164 influence of the HCl deposition parameters such as R<sub>HCl</sub> and N<sub>HCl</sub>, on the temporal change of  $J_{ev}$  and the mass accommodation coefficient  $\alpha$  during evaporation of a HCl-doped ice film 165 166 and its consequence on the lifetime of atmospheric ice particles contaminated by HCl. This 167 issue is key in relation to the importance of heterogeneous vs. homogeneous atmospheric 168 reactions at midlatitudes as has been pointed out in the past (Solomon et al., 1986; 1997).

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## 170 2. EXPERIMENTAL

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The emphasis of the present experiments was placed on the deposition of small amounts of HCl ranging in doses from 1 to 40 formal monolayers of HCl where a formal





monolayer of adsorbed HCl corresponded to a surface concentration of 2.5x10<sup>14</sup> molecule cm<sup>-</sup> 174 <sup>2</sup>. This is a consensus value obtained from several different experiments and is discussed at 175 176 length in Appendix 3 of reference 41. The apparatus as well as the methods used for 177 calibration and the HCl deposition procedure have been described in detail elsewhere (Delval 178 and Rossi, 2005). The experimental conditions are generally identical to the ones presented in 179 Delval and Rossi (2005) and the instrumental parameters are summarized in Table 1. The 180 only significant difference between the study of HNO<sub>3</sub>-doped ice and the present condensed phase investigation of HCl-doped ice lies in the mode of trace gas admission. HCl was 181 182 deposited by backfilling the reactor under stirred flow conditions with the inlet tubing used 183 for trace gas injection oriented towards one side of the Si-window of the cryostat set at 184 ambient temperature whereas HNO<sub>3</sub> was deposited by directed injection onto ice films 185 supported by the quartz crystal of the QCMB as referenced above. Evaporation experiments have been performed isothermally on samples in the temperature range 174-210 K under 186 187 dynamic pumping conditions, that is at maximum pumping speed (gate valve open) and at controlled temperature in order to prevent readsorption of HCl on the ice substrate. 188

First, an approximately 1.5 µm thick ice film was grown at 190 K on the quartz crystal 189 of the QCMB by deposition of bidistilled water vapor at a rate of  $1 \times 10^{17}$  molecule cm<sup>-2</sup> s<sup>-1</sup> 190 191 under static conditions. The H<sub>2</sub>O equilibrium vapor pressure agreed with published values 192 across the covered temperature range (Marti and Mauersberger, 1993; Mauersberger and 193 Krankowsky, 2003). Subsequently, the system was set to the desired temperature given in 194 Table 2 (second column from the left) and a metered amount of HCl was deposited under 195 stirred flow conditions. The rate of deposition of HCl, R<sub>HCl</sub>, as well as its time integral, namely the number of HCl molecules deposited on ice, N<sub>HCl</sub>, have been evaluated using the 196 method described in detail [41]. Typically,  $R_{HCl}$  ranges between  $8.0x10^{11}$  and  $4.2x10^{13}$ 197 molecule  $s^{-1}$  and N<sub>HCl</sub> between 1.0x10<sup>14</sup> and 5.4x10<sup>15</sup> molecules. The experimental conditions 198 199 of HCl-deposition as well as important experimental parameters are reported in Table 2. 200 Finally, the system was set to dynamic pumping conditions by opening the gate valve to the 201 pump.  $J_{ev}(H_2O)$  was measured isothermally using both the QCBM and residual gas MS. 202 Figure 1 illustrates a typical experimental protocol of the evaporation at 192 K of a HCl-203 doped ice film labelled as experiment 11 in Table 2 and performed as a multidiagnostic 204 experiment where both the gas- as well as the condensed phases are simultaneously 205 monitored.

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At t = 0, the system is set from stirred flow to dynamic pumping that starts the evaporation experiment. The continuous curve marked with the empty squares symbol in





Figure 1A corresponds to J<sup>QCM</sup><sub>ev</sub>, the evaporative flux of H<sub>2</sub>O calculated from the raw signal 208 209 at the output of the QCMB. The diamond symbol ( $\Diamond$ ) corresponds to  $J_{ev}^{18}$  evaluated from  $I_{ev}^{18}$ , the MS signal amplitude at m/e = 18.  $Int(J^{18}_{ev})$  marked by triangles in Figure 1A is the time 210 211 integral of J<sup>18</sup><sub>ev</sub> and corresponds to the total number of H<sub>2</sub>O molecules that have evaporated 212 from the ice film at t. D is the label at time t<sub>D</sub> at which J<sub>ev</sub>(H<sub>2</sub>O) decreased from its original 213 value corresponding to pure ice to 85% of its original value at t = 0, and  $d_D$  is the remaining 214 thickness of the ice film at t<sub>D</sub>. H<sub>b</sub> and H<sub>e</sub> in Figure 1B correspond to the time when HCl 215 evaporation begins and ceases to be observed, respectively, using gas phase residual mass 216 spectrometry (x symbols in Figure 1B) and are labeled t<sub>Hb</sub> and t<sub>He</sub>. The data have been treated 217 in analogy to HNO<sub>3</sub>-doped ice through the formalism given in Delval and Rossi (2005). Akin to HNO<sub>3</sub> the mass balance under dynamic pumping conditions between HCl deposited, 218 N<sup>dep</sup><sub>HCl</sub>, and HCl recovered during ice evaporation, N<sup>evap</sup><sub>HCl</sub>, agrees to within less than a factor 219 of 2. In the following N<sub>HCl</sub> will always refer to N<sup>dep</sup><sub>HCl</sub> derived from the measurement of HCl 220 221 at deposition because it refers to a directly measured quantity originating from a measured 222 pressure decrease in a given volume and time interval  $\Delta P/\Delta t$ ).

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#### 224 3. RESULTS

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The experimental data reported in Table 2 on the isothermal change of the evaporative 226 227 flux of water,  $J_{ev}(H_2O)$ , as a function of the average mole fraction of HCl,  $\chi_{HCl}$ , in the 228 remaining ice film during the evaporation process under dynamic conditions are presented in 229 Figure 2. Dynamic pumping conditions ensure the absence of any readsorption of  $H_2O$  vapor 230 during evaporation owing to the low H<sub>2</sub>O partial pressures in the reactor. The axes labelled 231 "b" and "e" correspond to the values of Jev(H2O) at the end of ice film deposition and after 232 desorption of most of the adsorbed HCl from the HCl-doped ice film at  $t_{He}$ , respectively, as 233 displayed in Figure 1B. The average mole fraction  $\chi_{HCl}$  of HCl in the remaining ice film as a 234 function of time is calculated according to Delval and Rossi (2005). The change in  $\chi_{HCI}$  owing 235 to  $H_2O$  evaporation is evaluated between  $t = t_D$  and  $t = t_{Hb}$  that corresponds to the time interval 236 when the number of adsorbed HCl molecules is constant as no release of HCl is observable in the gas phase at m/e = 36 before t<sub>Hb</sub>. Table 2 also displays the initial value of the HCl mole 237 238 fraction,  $\chi^0_{HCl}$ , calculated for the ice film just at the end of HCl deposition and marked by a 239 colored circle on the experimental trajectory of a color-coded evaporating ice film displayed 240 in Figure 2. The average mole fraction of HCl in the ice film,  $\gamma_{HCl}$ , increases owing to





241 evaporation of H<sub>2</sub>O from the ice film without loss of HCl such that the elapsed time increases

242 with  $\chi_{HCl}$  in Figure 2.

243 The beginning of an evaporation experiment after HCl doping (t = 0 in Figure 1) is 244 marked by a colored circle of a given experiment whose parameters are displayed in Table 2 245 and Figure 2. As pointed out above, at  $t = t_D J_{ev}(H_2O)$  has decreased to an arbitrarily chosen 246 value of 85% of its original value at  $t = t_0$  that corresponds to the beginning of the bold color-247 coded curve of a given experiment. Figure 2 essentially displays trajectories of evaporation experiments from t = 0 (colored circle) going to  $t_D$  and to finish at  $t_{Hb}$  between the two 248 249 limiting values for pure ice (color coded number of a given experiment on axis "b") and the 250 remaining ice film at the end of measurable HCl desorption t<sub>He</sub> (color-coded number of experiment on axis "e"). The trajectory of an experiment with reliable values of  $\chi_{HCl}$  from t = 251 252 0 (colored circle at  $\chi^0_{HCl}$ ) to t<sub>D</sub> and ending at t<sub>Hb</sub> is presented as a bold dashed-dotted and bold 253 smooth line from t = 0 to  $t_{Hb}$ , respectively, in order to emphasize the quantitative portion of 254 the experiment. Thinner (color-coded) dotted lines connect the phase prior to the start of 255 evaporation to the end of ice film deposition (colored circle on axis "b") and the post-phase of 256 evaporation starting at t<sub>Hb</sub> to t<sub>He</sub> in order to guide the eye of the reader to imagine a complete 257 evaporation cycle.

Two different data sets of the change of  $J_{ev}(H_2O)$  with  $\chi_{HCl}$  may be distinguished in 258 259 Figure 2. The first kind of data set corresponds to the curves describing  $J_{ev}$  for experiments 1, 2, 9 and 11 and is called dataset A. These traces present a slow continuous decrease of 260 261  $J_{ev}(H_2O)$  as  $\chi_{HC1}$  increases during  $H_2O$  evaporation. The second type of dataset shows an 262 initial plateau of  $J_{ev}(H_2O)$  with increasing  $\chi_{HCI}$  starting at the value of pure ice evaporation 263 followed by a sudden decrease of  $J_{ev}(H_2O)$  and is found for experiments 3, 4, 7 and 8 which 264 we call dataset B. Akin to HNO<sub>3</sub>, we have evaluated the impact of the HCl deposition protocol on the evaporation range parameter, r<sup>b/e</sup>, which is the ratio between the evaporative 265 flux of H<sub>2</sub>O at the beginning of ice evaporation,  $J_{ev}^{b}(H_2O)$  reported on the left axis "b" in 266 Figure 2, and  $J_{ev}(H_2O)$  close to the end of the desorption of HCl,  $J_{ev}^e(H_2O)$ , at  $t = t_{He}$  (the right 267 axis "e" in Figure 2). It describes the factor by which Jev(H2O) decreases within the limits of 268 "b" and "e". The impact of both the rate of deposition of HCl on ice, R<sub>HCl</sub>, and its time 269 270 integral corresponding to the dose of deposited HCl, N<sub>HCl</sub>, are presented in Figures 3 and 271 Figure A1 (Appendix), respectively.

It appears from these Figures that we have not succeeded to find a simple experimental parameter that controls  $J_{ev}(H_2O)$  either with elapsed time or amount of adsorbed HCl expressed as the time dependence of  $\chi_{HCl}$ . Instead, the data may roughly be classified along





the two cases presented above, namely datasets A and B. The distinction between both data sets seems to be the rate of change (slope) of  $J_{ev}(H_2O)$  within a narrow range of  $\chi_{HCl}$ . Indeed, the available number of experiments clearly shows two distinct and limiting cases whereas the search for other controlling parameters such as  $R_{HCl}$ ,  $N_{HCl}$  and the temperature of deposition ( $T_{ice}$ ) for dataset A failed akin to the HNO<sub>3</sub> study (Delval and Rossi, 2005).

280 One may take note for instance of the low value of  $\chi_{HCl}$  at 210 K for experiment 9 281 where the conditions of deposition are similar to experiments 1 and 2, yet, its respective values of  $r^{b/e}$  differ significantly from experiment 9 (Figure 3). In contrast, for dataset B the 282  $r^{b/e}$  values are similar for the whole set and range from 20 to 27.2 staying within a fairly 283 narrow band. Moreover, they seem to be independent of R<sub>HCl</sub> and N<sub>HCl</sub> as for data set A. In 284 contrast, the r<sup>b/e</sup> values for dataset A seem widely scattered over the explored parameter space. 285 We have also investigated the impact of the deposition protocol on d<sub>D</sub>, which is the thickness 286 of ice that is affected by the presence of HCl, namely the remaining thickness of ice whose 287 288 Jev(H2O) value has decreased to 85% of Jev(H2O) of pure ice. The results on dD as a function of R<sub>HCl</sub> and N<sup>dep</sup><sub>HCl</sub> are presented in Figures 4 and A2 (Appendix), respectively. Taking the 289 results of Figures 3, 4, A1 and A2 together we arrive at the following two conclusions: 290

(1) T<sub>ice</sub>, R<sub>HCl</sub> and N<sup>dep</sup><sub>HCl</sub> are not controlling parameters or predictors for J<sub>ev</sub>(H<sub>2</sub>O) of either set.
(2) The evaporation range parameters r<sup>b/e</sup> and d<sub>D</sub> are not characterizing set A. In contrast, for
dataset B, r<sup>b/e</sup> and d<sub>D</sub> values fall into a narrow range with values varying from 460.7 to 636.0
nm compared to the original ice thickness d<sub>0</sub> of 1'500 nm or so (exact numbers in Table 2).

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#### 296 4. DISCUSSION

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298 Figure 1 displays the evaporation history of sample 11 whose deposition parameters are listed in Table 1. The initial average mole fraction  $\chi_{HCl}^{0}$  of HCl, once deposition on the 1.44 299 300 um thick ice film under stirred flow reactor conditions is terminated, has been estimated from 301 the total number of H<sub>2</sub>O molecules contained in the ice film and the measured number of deposited HCl molecules, N<sup>dep</sup><sub>HCl</sub>, for experiment 11 (Table 2). Table 2 and Figure 1 reveal 302 that for approximately  $2.2 \times 10^{18}$  H<sub>2</sub>O molecules in the film and  $5.4 \times 10^{14}$  molecules of 303 deposited HCl, we obtain  $\chi_{HCl}^{0} = 2.7 \times 10^{-4}$ . This HCl mole fraction represents an average 304 305 value that takes into account all H<sub>2</sub>O molecules contained in the ice film whereas in reality 306 there will be a HCl gradient across the ice film as has been observed in the case of the 307 HNO<sub>3</sub>/ice system (Delval and Rossi, 2005).





308 After the HCl deposition process on the typically 1.5 µm thick ice film the gate valve is 309 opened in order to initiate the isothermal evaporation experiment under dynamic pumping 310 conditions. Initially, H<sub>2</sub>O evaporates at fluxes J<sub>ev</sub>(H<sub>2</sub>O) that are characteristic of pure ice 311 measured previously (Delval and Rossi, 2004; Pratte et al., 2006). These initial values  $J_{ev}^{b}(H_2O)$  are displayed on the left-hand "b" (= "beginning") axis in Figure 2. As the 312 evaporation proceeds  $J_{ev}(H_2O)$  slightly decreases with time as displayed in Figure 1A to the 313 314 arbitrarily chosen point where  $J_{ev}(H_2O)$  has decreased to 85% of the initial pure ice value at 315 which point the remaining ice thickness  $d_D$  has decreased by approximately one third to 771.7 316 nm remaining ice thickness as displayed in Figure 1B and Table 2. Further evaporation of 317 H<sub>2</sub>O leads to a continuous decrease of  $J_{ev}(H_2O)$  at a corresponding increase of  $\gamma_{HCl}$  up to point 318  $H_b$  defined above ("Halogen beginning") at  $t_{Hb}$  (Figure 1B) where HCl starts to desorb from 319 the ice film as monitored using the residual MS signal at m/e = 36.

320 For t < t<sub>Hb</sub>,  $\chi_{HCl}$  is given by the number of originally deposited HCl molecules that 321 remain adsorbed on the ice film up to t<sub>Hb</sub> and the remaining H<sub>2</sub>O molecules in the film. In contrast, for  $t > t_{Hb}$  the composition of the remaining ice film must be determined by taking 322 323 into account the loss by evaporation of both H<sub>2</sub>O and HCl. The present experimental 324 configuration is not adapted to quantitatively measure HCl loss. Therefore, we have chosen to 325 display the temporal development of  $J_{ev}(H_2O)$  for  $t < t_{Hb}$  in Figure 2 as a function of the 326 average value of the HCl mole fraction  $\chi_{HCl}$ . However, the value of  $J_{ev}(H_2O)$  at  $t = t_{He}$  where 327 most of the HCl has desorbed from the ice film is plotted on the right axis labelled "e" (= 328 "end") as  $J_{ev}^{e}(H_2O)$  in Figure 2 in order to provide a limit for the minimum value of the 329 evaporation rate  $J_{ev}(H_2O)$  at an ice film thickness  $d_{He}$  of approximately 80 ± 10 nm as 330 displayed in Figure 1B. We have observed in the past that  $J_{ev}(H_2O)$  for a pure ice film of 331 approximate thickness of 80 nm or less also slows down, presumably owing to island 332 formation at the very end of pure thin ice film evaporation [41]. Therefore, results are becoming more difficult to interpret such that we halted the experiment at  $t_{He}$ . The ratio  $r^{b/e} =$ 333 J<sub>ev</sub><sup>b</sup>(H<sub>2</sub>O)/J<sub>ev</sub><sup>e</sup>(H<sub>2</sub>O) is displayed in Table 2 and is an operational evaporation range parameter 334 335 that estimates the extent of decrease of  $J_{ev}(H_2O)$  for a thick HCl-doped ice film of  $\mu m$  size 336 down to thicknesses of approximately 80 nm.

At the start of the evaporation experiment the equilibrium vapor pressure of H<sub>2</sub>O, P<sub>eq</sub>(H<sub>2</sub>O), is that of pure ice (Delval et al., 2003; Delval and Rossi, 2004; Pratte et al., 2006) owing to the small values of  $\chi_{HCl}^{0}$ . Raoult's Law applies to such small values of  $\chi_{HCl}$  but leads to unmeasurably small deviations from the observed vapor pressure of H<sub>2</sub>O which is that of pure ice. In fact, we have never observed an equilibrium vapor pressure that did not





342 correspond to pure ice in the course of the present work that seems to be the consequence of 343 the small average mole fractions of HCl in the H<sub>2</sub>O/HCl system. This value of P<sub>eq</sub>(H<sub>2</sub>O) is observed throughout the evaporation up to  $t_{He}$  as the film is apparently sufficiently H<sub>2</sub>O-rich 344 345 to support an equilibrium vapour pressure characteristic of pure ice. In view of the decreasing values of  $J_{ev}(H_2O)$  displayed in Figure 2 the equilibrium vapour pressure of pure ice can only 346 347 be maintained if the condensation rate coefficient kc for H2O adsorption decreases to the same 348 extent as  $J_{ev}(H_2O)$  in agreement with previous work (Delval et al., 2003; Delval and Rossi, 349 2004; Pratte et al., 2006) and the concept of microscopic reversibility.

350 Figure 1A displays both the QCMB signal (□) as well as the corresponding MS signal 351 for evaporating H<sub>2</sub>O at m/e = 18 ( $\diamond$ ). Akin to the HNO<sub>3</sub>/H<sub>2</sub>O system studied previously 352 (Delval and Rossi, 2005) we obtain a perfect match between the two signals for  $t < t_D$  whereas 353 for  $t > t_D$  there is a significant discrepancy, especially at t > 300 s amounting to typically less than a factor of two. Such a disagreement has been noted before for  $HNO_3/H_2O_3$  albeit to a 354 355 larger extent. The reason for this behaviour of the QCMB signal has not been studied in detail but may well lie in a structural rearrangement of the condensed phase during 356 357 evaporation that will lead to a change in the calibration factor  $C_f$  defined in Table 1 and in 358 Delval and Rossi (2005). In view of the straightforward interpretation of the calibrated MS 359 signal at m/e = 18 we have used it for the measurement of  $J_{ev}(H_2O)$  at  $t > t_D$  akin to the previous study on HNO<sub>3</sub>/H<sub>2</sub>O. 360

361 The accuracy with which both t<sub>Hb</sub> and t<sub>He</sub> can be determined depends on the temporal 362 change of the background MS signal for HCl at m/e = 36 displayed in Figure 1B following 363 the dosing of the thin ice film under stirred flow conditions. Figure 1B displays the MS signal at m/e = 36 as a function of time just before the start of HCl desorption at  $t_{Hb}$  that is signalled 364 by an increase in the MS intensity whereas  $t_{He}$  corresponds to the return of the HCl signal to 365 366 the decaying HCl background in comparison to a reference experiment in which the HCl 367 background was monitored as a function of time following the admission of the same HCl 368 dose in the absence of an ice film. We estimate that  $t_{Hb}$  is determined to  $\pm 10$  s whereas  $t_{He}$ may only be estimated to  $\pm$  100 s by virtue of the vanishing intensity of the HCl MS signal 369 370 compared to its changing background.

371 Previous work has established that the rate of deposition of HCl,  $R_{HCl}$ , in the range 372  $1 \times 10^{13}$  to  $5 \times 10^{13}$  molecule s<sup>-1</sup> for the 0.78 cm<sup>2</sup> surface area of the Si-window leads to the 373 formation of a crystalline HCl hydrate, HCl•xH<sub>2</sub>O, whereas values outside of this range 374 seemed to favor the formation of an amorphous HCl/H<sub>2</sub>O mixture (Delval et al., 2003). The 375 exact nature of this undoubtedly crystalline solid is still unknown. However, IR spectroscopic





376 work on hydroxonium salts of the type  $H_3O^+X^-$  suggests that the v<sub>1</sub> and v<sub>3</sub> peak positions of 377 the symmetric and antisymmetric O-H stretch vibrations must correspond to a molecular 378 structure in which the distance between the cation and anion is unusually large (Desbat and 379 Huong, 1975; Iannarelli and Rossi, 2016). Recent work has shown that the presence of HCl 380 hexahydrate (HCl•6H<sub>2</sub>O) under the present experimental conditions could be safely excluded, 381 however, the FTIR absorption spectrum clearly shows the presence of dissociated HCl within 382 the ice film (Iannarelli and Rossi, 2014). Akin to HCl•6H<sub>2</sub>O that is known to nucleate with difficulty, crystallization of this unknown HCl hydrate seems to occur only under specific 383 conditions of temperature and/or HCl deposition. Owing to the quantitative control of HCl 384 385 deposition on the ice film in this work we infer the presence of at least two forms of HCl hydrates in the temperature range chosen in analogy to previous work (Delval et al., 2003). 386

387 We clearly point out that the present work has been performed without simultaneous spectroscopic control of the HCl/ice deposit that would have allowed the identification and/or 388 389 quantification of the molecular composition of the condensate. Because we lack a 390 spectroscopic probe for the ice film deposited on the QCMB in the present work we are 391 seeking a correlation between the type of HCl/H<sub>2</sub>O deposit, either crystalline or amorphous, 392 and the relevant HCl deposition parameters. Previous work has revealed a distinctly different 393 temporal dependence of  $J_{ev}(H_2O)$  between the crystalline and amorphous HCl hydrates with 394 the extent of  $H_2O$  evaporation from the film, both at low (Delval et al., 2003) and high 395 temporal resolution (Iannarelli and Rossi, 2014).

396 Datasets A and B have been characterized above in terms of a difference in the temporal 397 dependence of  $J_{ev}(H_2O)$  as a function of  $\chi_{HCl}$ . Taking one example of each set Figure 2 398 reveals a distinct difference between experiment 7 (set B) and 11 (set A) performed at T =399 195 and 192 K, respectively, despite comparable HCl deposition parameters (Table 2). At t >  $t_D J_{ev}(H_2O)$  for experiment 7 decreases at once with  $\chi_{HCI}$  in contrast to experiment 11 whose 400 401  $J_{ev}(H_2O)$  value gradually starts to decrease at roughly the same value of  $\chi_{HCl}$  as experiment 7. 402 In addition, in both cases the extent of the decrease of  $J_{ev}(H_2O)$  is roughly equal between  $t_D$ 403 and  $t_{Hb}$ . Set B data are in marked contrast to set A independent of the magnitude of  $\chi_{HCI}$ 404 which is highlighted by a comparison of experiment 11 (set A) and 4 (set B) at 192 and 190 405 K, respectively. Both the abrupt decrease of  $J_{ev}(H_2O)$  at t<sub>D</sub> for set B as well as the gradual 406 decline for set A occur before HCl starts to evaporate from the sample at  $t_{Hb}$  and appear to be 407 independent of  $\chi_{HCl}$  within the range explored in the present work. However, we admit that we 408 do not have any information on the HCl profile in the remaining ice film.





409 If we consider the mean value  $\langle d_D \rangle$  for data set B (Figures 4 and A2) we find 549.0 ± 120.0 nm compared to the 1'500 nm or so original ice thickness which corresponds to 410  $8.5 \times 10^{17}$  molecules of H<sub>2</sub>O spread out over 0.50 cm<sup>2</sup>. These H<sub>2</sub>O molecules are impacted by 411 the presence of HCl to some extent because Jev(H2O) is slowed down significantly compared 412 413 to pure ice. Previous results (Delval et al., 2003) on the deposition of HCl on ice under 414 conditions where the presence of an as yet unidentified crystalline hydrate HCl•xH<sub>2</sub>O was 415 confirmed by FTIR absorption led to the conclusion that on average the amount of "trapped" H<sub>2</sub>O within d<sub>D</sub> corresponded to  $1.2 \times 10^{18}$  H<sub>2</sub>O starting with an original 1 µm thick ice film that 416 was subsequently doped with HCl. This quantity of H<sub>2</sub>O, when scaled from the 0.78 cm<sup>2</sup> area 417 of the Si-window used for FTIR absorption to the area of 0.5 cm<sup>2</sup> of the OCMB leads to 418  $7.7 \times 10^{17}$  H<sub>2</sub>O that is in satisfactory agreement with the present measurement of d<sub>D</sub> or  $8.5 \times 10^{17}$ 419  $H_2O$  in the present work. We may add that the previous value of  $1.2 \times 10^{18} H_2O$  from the work 420 421 of Delval et al. (2003) corresponding to d<sub>D</sub> obtained in that work has been derived using 422 HeNe interferometry which is a crude method for measuring the film thickness.

423 Specifically, considering the low value of d<sub>D</sub> of experiments 1 and 10 (Table 2, Figure 424 4) we may define the behaviour of these condensates as "ice-like" because roughly 80% of the 425 ice sample of roughly 1.5  $\mu$ m thickness has evaporated at J<sub>ev</sub>(H<sub>2</sub>O) of pure H<sub>2</sub>O ice before it slows down. This decrease of Jev(H2O) is a kinetic effect and acts on both the rate of 426 427 evaporation as well as on the mass accommodation coefficient, the ratio of which remains 428 constant because the characteristic vapor pressure of pure ice is maintained until  $t = t_{He}$  when 429 the sample runs out of H<sub>2</sub>O and HCl. For sample 1 this conclusion is not too surprising owing 430 to its extremely low HCl dose of 0.8 formal HCl monolayers. Sample 10 in comparison with the other members of data set A allows us to conclude that  $d_D$  is proportional to  $T_{ice}$  for data 431 432 set A. Low temperatures prevent rapid diffusion of HCl into the bulk of the ice film which 433 leaves the majority of the total mass of the thin film deposited void of any HCl. Therefore, a 434 large fraction of the total mass of the thin film deposit evaporates at values of J<sub>ev</sub>(H<sub>2</sub>O) 435 characteristic of pure ice before it decreases to lower values when the presence of HCl slows 436 down J<sub>ev</sub>(H<sub>2</sub>O). Although our experiment does not reveal the location of the thin layer of HClcontaminated ice, plausibility suggests that it is located on top of the ice film at the gas-437 438 condensed interface. The corollary of this is that it is impossible to "cap" a pure ice sample 439 with a thin layer of an atmospheric condensable gas of lower vapor pressure in the hope to 440 lower the vapour pressure of the condensate or slow down H<sub>2</sub>O evaporation. This capping has 441 been attempted many times, and examples abound. However, all attempts to lower the ice





vapor pressure of the condensate using low amounts of polar contaminants of ice, such as
HNO<sub>3</sub>, HCl or HBr have proven futile to date (Biermann et al., 1998).

The other members of data set A are examples (experiments 2, 9, 11) with high values of  $d_D$  at higher temperatures and higher HCl doses (Table 2). Because of higher presumed interfacial HCl concentrations these samples experience a decrease in  $J_{ev}(H_2O)$  owing to rapid diffusion of HCl into ice that affects the kinetics of evaporation to some depths of the ice film corresponding to higher values of  $d_D$ . Both high HCl doses and high temperatures favor HCl contamination of deeper layers of the HCl film, hence high values of  $d_D$ .

450 Tentatively, we assign a crystalline, yet unknown molecular structure and stoichiometry 451 to samples A in contrast to samples of dataset B that we identify with an amorphous structure 452 in terms of a liquid HCl/H<sub>2</sub>O mixture of variable composition. The main argument in favour 453 of this assignment comes from recent kinetic work performed by Iannarelli and Rossi (2016a) who show that both  $J_{ev}(H_2O)$  as well as the corresponding mass accommodation coefficient or 454 the adsorption rate coefficient for H<sub>2</sub>O adsorption is highly scattered for crystalline HCl 455 hexahydrate whereas the amorphous mixture shows a significantly smaller scatter of the 456 457 experimental and thermodynamic values (Iannarelli and Rossi, 2014). Figure A3 and A4 in 458 the Appendix show this substantial difference in experimental scatter for the amorphous HCl/H<sub>2</sub>O mixture (Figure A3) compared to crystalline HCl hexahydrate (Figure A4). 459

Figure 3 displays the range parameter  $r^{b/e}$  as a function of  $R_{HCl}$  for all data displayed in Table 2. It is noteworthy that  $r^{b/e}$  is in the range 20 to 27 for set B experiments 3, 4, 7 and 8 compared to set A data that seem to be scattered throughout the range. Members of data set B show a common average range for both  $d_D$  and  $r^{b/e}$  which is the reason we tentatively assign these structures to amorphous liquid mixtures of high viscosity at the prevailing temperatures.

465 In conclusion, we take the simultaneous occurrence of the restricted range of the measured remaining thickness of ice  $d_D = 549.0 \pm 120.0$  nm together with a similarly 466 restricted range of r<sup>b/e</sup> between 20 and 27 as well as the substantial overlap in R<sub>HCl</sub> between 467 the present and previous work (Delval et al., 2003) as an indication that set B evaporation 468 experiments imply the presence of an amorphous HCl/H<sub>2</sub>O mixture. In contrast, the scatter of 469 the set A data across the range of r<sup>b/e</sup> and d<sub>D</sub> values suggests the presence of an as yet 470 471 unidentified crystalline HCl hydrate. It remains to be seen whether or not the published FTIR 472 absorption spectrum published in Delval et al (2003) turns out to be identical to the 473 hypothetical crystalline HCl hydrate invoked as condensate in set A molecules, similar HCl 474 deposition parameters notwithstanding. This proposal awaits further confirmation from FTIR 475 spectroscopic work that will be combined in the future with the QCMB measurement. At this





476 point we reiterate our earlier statement that  $T_{ice}$ ,  $R_{HCl}$ ,  $N^{dep}_{HCl}$  do apparently not control 477  $J_{ev}(H_2O)$  of both sets A and B.

478

#### 479 5. ATMOSPHERIC IMPLICATIONS

480

481 The evaporation range parameter  $r^{b/e}$  may be used to quantitatively evaluate the 482 evaporative lifetime extension of thin ice films under conditions of H<sub>2</sub>O vapor subsaturation. 483 In the interest of applying the data of the present work to atmospheric conditions we make the 484 assumption that typical atmospheric Cirrus cloud particles of several  $\mu$ m diameter may be 485 approximated by macroscopic thin films used to obtain the present data. The time t<sub>ev</sub> in 486 seconds to complete evaporation of an ice particle of radius r at a given relative humidity (rh) 487 is given in equation 1 (Chiesa and Rossi, 2013; Iannarelli and Rossi, 2016a):

488 
$$t_{ev} = \frac{\left(\frac{\rho N_L}{M}\right)^{2/3} \left(\frac{r}{a}\right)}{J_{ev} (1 - rh)}$$
(1)

where  $\rho$  is the density of ice (0.916 and 0.925 g cm<sup>-3</sup> at 273 and 173 K, respectively), 489  $M = 18 \text{ g mol}^{-1}$  of H<sub>2</sub>O, r and a are the ice particle radius and the distance between two 490 molecular layers of H<sub>2</sub>O(ice) in ice, respectively (Iannarelli and Rossi, 2016a). Equation 1 is 491 492 based on a simple layer-by-layer evaporation model of H<sub>2</sub>O(ice) from a spherical ice particle 493 following a zero-order rate law for J<sub>ev</sub> or a first order rate law for its inverse, namely H<sub>2</sub>O adsorption or condensation. For a 10 µm diameter ice particle approximated by thin film 494 experiment 1 (Table 2) at rh = 80%, T = 192 K,  $J_{ev} = 3x10^{16}$  molecule s<sup>-1</sup> cm<sup>-2</sup> (Petrenko and 495 Whitworth, 1999) and a =  $4 \times 10^{-8}$  cm we obtain t<sub>ev</sub> = 2050 s or 34 min. This is the value for a 496 497 pure ice particle as  $J_{ev}(H_2O)$  for pure ice has been used at the outset of the evaporation experiment and is a lower limit to the true evaporation time owing to the competition of mass 498 transfer and heterogeneous chemistry (Seinfeld and Pandis, 1998). Using  $r^{b/e} = 43$  for 499 experiment 1  $t_{ev}$  is calculated to be 15 minutes and 24 hours for a 100 nm and 10  $\mu$ m diameter 500 particle, respectively, whereas the evaporative lifetime of an analogous pure ice particle 501 502 would be only 21 s for the 100 nm diameter pure ice particle. Cirrus ice particles are 503 frequently in the lower tens of µm size range resulting in a longer evaporation time 504 considering that the simple evaporation model scales linearly with the radius of the ice 505 particle. In conclusion we may state that owing to the lifetime extension of ice particles 506 contaminated by HCl, HNO<sub>3</sub> or other volatile atmospheric trace gases such as HOCl, HOBr or 507 HONO small particles may have a chance to survive subsaturated regions of the atmosphere





so as to function as cloud condensation or ice nuclei for the following cloud cycle (Delval and
Rossi, 2004; 2005; Pratte et al., 2006).

We would like to stress that the variable r<sup>b/e</sup> factor displayed in Table 2 leads to a 510 significant increase in the evaporative lifetime of a contaminated ice particle and amounts to a 511 512 kinetic effect that does not affect the equilibrium vapor pressure of the ice particle in 513 question: it is that of pure ice from the start of the evaporation experiment to  $t = t_{He}$  and 514 therefore affects both the rate of evaporation and accommodation equally. However, in cases 515 the sample has lost most of its mass the vapor pressure decreases and becomes somewhat 516 uncertain. In the present case the above statement is correct for  $t = t_{Hb}$ , that is before halogen evaporation. Of note is the fact, that the accommodation coefficient  $\alpha$  is frequently less than 517 518 unity, in contrast to what is often assumed, which will lower the rate of evaporation for pure 519 ice, hence increase the evaporative lifetime of pure ice particles for  $T \ge 180$  K as proposed in 520 previous work (Delval and Rossi, 2004; 2005; Pratte et al., 2006).

521 As a token example of potential atmospheric importance of the measured evaporative 522 lifetimes of ice particles laced with condensable atmospheric trace gases we may take the 523 formation, persistence and evaporation of contrails and Cirrus clouds in the UT/LS. These are 524 ice clouds forming on non-volatile ice nuclei at the corresponding temperature and relative 525 humidity conditions and also frequently serve as reaction sites for heterogeneous atmospheric 526 reactions in connection with ozone depletion and chlorine activation chemistry in the LS. 527 Under certain conditions, Schumann and coworkers used the concept of the increase of the 528 evaporative lifetimes of contaminated ice particles in aviation contrails occurring mostly in 529 the UT, but sometimes also in the LS, in order to explain the persistence of ice clouds below 530 ice saturation conditions up to a certain time duration. Ice clouds have a significant radiative 531 forcing effect that is of interest in evaluating the climate forcing of high-flying aircraft in 532 future aviation scenarios (Lewellen, 2014; Schumann et al, 2017; 2017a). However, the 533 results of the present work show that the rate of evaporation of ice films doped with small 534 amounts of acidic trace gases significantly slows down in a complex manner over the 535 evaporation history of the film or particle, and that the application of equation (1) to 536 atmospheric situations should be carried out with caution.

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# 542 CONCLUSIONS

544	Despite the scatter of the values of $r^{b/e}$ and $d_{\rm D}$ in dataset A displayed in Figures 3 and 4
545	and the apparent lack of influence of the deposition parameters $(T_{ice},\ R_{HCl},\ N^{dep}_{HCl})$ on
546	$J_{ev}(H_2O)$ we may state several key points from the present work:

- (a) We observe two types of behaviour, both complex, as far as the temporal change of
   J<sub>ev</sub>(H<sub>2</sub>O) with on-going evaporation of H<sub>2</sub>O from a HCl/H<sub>2</sub>O condensate is
   concerned. We have named it sets A and B that represent limiting behaviour as not
   all performed experiments fit into this scheme.
- (b) At low temperature or low dose of deposited HCl ( $N^{dep}_{HCl}$ ) set A samples, especially samples 1 and 10, reveal an "ice-like" behaviour that corresponds to a low value of d<sub>D</sub>. This means that the HCl/H<sub>2</sub>O condensate evaporates a large fraction of the sample thickness at a value of J<sub>ev</sub>(H<sub>2</sub>O) characteristic of pure ice before slowing down at increasing mole fraction of HCl upon H<sub>2</sub>O evaporation. This corresponds to a two-phase system consisting of a major ice-like and a minor HCl/H<sub>2</sub>O phase having both significantly different values of J<sub>ev</sub>(H<sub>2</sub>O).
- 558(c) High values of  $d_D$  are observed at high  $T_{ice}$  or  $N^{dep}_{HCl}$  values for set A samples. This559means that the sample evaporates  $H_2O$  at  $J_{ev}(H_2O)$  characteristic of pure ice for a560relatively short time of its evaporation history because the quantity of HCl is561sufficient to decrease  $J_{ev}(H_2O)$  already at high values of  $d_D$  by rapidly diffusing to562deeper layers of the ice film. An equivalent way of expressing the point would be to563state that  $d_D$  which is an indicator of the total mass of the ice film, is proportional to564 $T_{ice}$  for Set A.
- 565 (d) Set A samples generally show scattered values of both  $d_D$  and  $r^{b/e}$  values that we 566 attribute to the existence of a two-phase binary system, namely a pure ice and a 567 crystalline HCl hydrate phase of unknown stoichiometry HCl•xH<sub>2</sub>O. At first the 568 pure ice phase starts to evaporate as a whole for a fairly long time at characteristic 569 values of  $J_{ev}(H_2O)$  until the pure ice phase has disappeared, followed by the 570 crystalline HCl/H<sub>2</sub>O phase at a lower rate of  $J_{ev}(H_2O)$  to attain the characteristic 571 value for the evaporation of the crystalline HCl•xH<sub>2</sub>O phase.





576		manifests itself as a continuous, yet gradual decrease of $J_{ev}(\mathrm{H_2O})$ with increasing
577		$\chi_{HCl},$ in distinct contrast to Set A where $J_{ev}(H_2O)$ values are those of pure ice until
578		the ice phase has completely evaporated following which a gradual decline of
579		$J_{ev}(H_2O)$ ensues when the crystalline HCl hydrate starts to decompose.
580	(f)	It must be recalled that the vapour pressure of $\mathrm{H_{2}O}$ remained that of pure ice during
581		most of the thickness of the $\mathrm{H_{2}O/HCl}$ condensate down to approximately 80 nm at
582		which point we halted the evaporation experiment. This result is expected based on
583		Raoult's law owing to the small average HCl mole fractions in doped ice used in the
584		present work: It would make the decrease of the H2O saturation vapour pressure
585		unmeasurably small. The present results therefore primarily address the kinetics of
586		$\mathrm{H_{2}O}$ evaporation which changes with the total mass of the thin film condensate and
587		the concomitant increase in HCl concentration and/or mole fraction.
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589	References
590	
591	Abbatt, J. P. D. Interactions of atmospheric trace gases with ice surfaces: Adsorption and
592	reactions, Chem. Rev. 2003, 103, 4783-4800.
593	Banham, S. F.; Horn, A. B.; Koch, T. G.; Sodeau, J. R. Ionisation and solvation of
594	stratospherically relevant molecules on ice films, Faraday Discuss. 1995, 100, 321-
595	332.
596	Biermann, U.; Crowley, J. N.; Huthwelker, T.; Moortgat, G. K.; Crutzen, P. J.; Peter, T. FTIR
597	studies on lifetime prolongation of stratospheric ice particles due to NAT coating,
598	Geophys. Res. Lett. 1998, 25, 3939-3942.
599	Bolton, K.; Petterson, J. B. C. Ice-Catalyzed Ionization of Hydrochloric Acid, J. Amer. Chem.
600	Soc. 2001, 123, 7360-7363.
601	Bournel, F.; Mangeney, C.; Tronc, M.; Laffon, C.; Parent, P. Acidity of hydrogen chloride at
602	the surface of low-temperature .40-150 K. water-ice films, Phys. Rev. B 2002, 65,
603	201404-1 to 201404-4.
604	Buch, V.; Sadlej, J.; Aytemiz-Uras, N.; Devlin, J. P. Ice-Catalyzed Ionization of Hydrochloric
605	Acid J. Phys. Chem. A 2002, 106, 9374-9389.
606	Chiesa, S.; Rossi, M. J. $HCl \cdot 6H_2O$ is metastable - IR spectroscopy, phase transitions and
607	kinetic/thermodynamic properties in the range 170-200 K, Atm. Chem. Phys. 2013,
608	<i>13</i> , 11905-11923.
609	Chu, L. T.; Leu, MT.; Keyser, L. F. Uptake of HCI in Water Ice and Nitric Acid Ice Films,
610	J. Phys. Chem. <b>1993</b> , <i>97</i> , 7779-7785.
611	Delval, C.; Flückiger, B., Rossi, M. J. The Rate of Water Vapor Evaporation from Ice
612	Substrates in the presence of HCl and HBr: Implications for the Lifetime of
613	atmospheric Ice Particles, Atmos. Chem. Phys. 2003, 3, 1131-1145.
614	Delval, C.; Rossi, M. J. The kinetics of condensation and evaporation of H <sub>2</sub> O from pure ice in
615	the range 173 to 223 K: A quartz crystal microbalance study, Phys. Chem. Chem.
616	Phys. <b>2004</b> , <i>6</i> , 4665-4676.
617	Delval, C.; Rossi, M. J. The influence of monolayer amounts of HNO3 on the evaporation rate
618	of H <sub>2</sub> O over ice at $179 \le T/K \le 208$ : A quartz crystal microbalance study, J. Phys.
619	Chem. A <b>2005</b> , <i>109</i> , 7151-7165.
620	Delzeit, L.; Rowland, B.; Devlin, J. P. Infrared Spectra of HCl Complexed/Ionized in
621	Amorphous Hydrates and at Ice Surfaces in the 15-90 K Range, J. Phys. Chem. 1993,
622	97, 10312-10318.





- 623 Delzeit, L.; Rowland, B.; Devlin, J. P. Ice Surface Reactions with Acids and Bases, J. Phys.
- 624 Chem. **1993a**, 97, 10312-10318.
- Delzeit, L.; Powell, K.; Uras, N.; Devlin, J. P. Ice Surface Reactions with Acids and Bases, J.
  Phys. Chem. B 1997, *101*, 2327-2332.
- 627 Desbat, B.; Huong, P. V. Spectres i.r. et Raman des sels d'hydroxonium  $H_3O^+CI^-$ ,  $H_3O^+Br^-$  et 628  $H_3O^+SbCl_6^-$ , Spectrochim. Acta, Part A **1975**, *31*, 1109-1114.
- Devlin, J. P.; Uras, N.; Sadlej, J.; Buch, V. Discrete stages in the solvation and ionization of
  hydrogen chloride adsorbed on ice particles, Nature 2002, *414*, 269-271.
- Donsig, H. A.; Vickerman, J. C. Dynamic and static secondary ion mass spectrometry studies
  of the solvation of HCl by ice, J. Chem. Soc. Faraday Trans. 1997, 93, 2755-2761.
- Ferriso, C. G.; Hornig, D. F. Infrared Spectra of Oxonium Halides and the Structure of the
  Oxonium Ion, J. Chem. Phys. 1955, 23, 1464-1468.
- Flückiger, B.; Thielmann, A.; Gutzwiller, L.; Rossi, M. J. Real-Time Kinetics and
  Thermochemistry of the Uptake of HCl, HBr and HI on Water Ice in the Temperature
  Range 190 to 210 K, Ber. Bunsenges. Phys. Chem. **1998**, *102*, 915-928.
- Flückiger, B.; Rossi, M. J. Common Precursor-mediated Reaction Mechanism for the
  Heterogeneous Interaction of D<sub>2</sub>O, HCl, HBr and HOBr on Ice at low Temperatures, J.
  Phys. Chem. A 2003, 107, 4103-4115.
- Gertner, B. J.; Hynes, J. T. Molecular Dynamics Simulation of Hydrochloric Acid Ionization
  at the Surface of Stratospheric Ice, Science 1996, 271, 1563-1566.
- Gilbert, A. S.; Sheppard, N. Infra-red Spectra of the Hydrates of Hydrogen Chloride and
  Hydrogen Bromide Absorption Bands of the H<sub>5</sub>O<sub>2</sub><sup>+</sup> Species, J. Chem. Soc. Faraday
  Trans. 1973, 69, 1628-1642.
- 646 Graedel, T. E.; Keene, W. C. Tropospheric budget of reactive chlorine, Global
  647 Biogeochemical Cycles 1995, *9*, 47-77.
- Graham, J. D.; Roberts, J. T. Interaction of HCl with crystalline and amorphous ice:
  implications for the mechanisms of ice-catalyzed reactions, Geophys. Res. Lett. 1995,
  22, 251-254.
- Graham, J. D.; Roberts, J. T. Formation of HCl•6H<sub>2</sub>O from ice and HCl under ultrahigh
  vacuum, Chemom. Intell. Lab. Systems 1997, *37*, 139-148.
- Hanson, D.R.; Ravishankara, A.R. Investigation of the Reactive and Nonreactive Processes
  Involving ClONO<sub>2</sub> and HCl on Water and Nitric Acid Doped Ice, J. Phys. Chem.
- 655 **1992**, *96*, 2682-2691.





656	Houghton, J. T. et al.; eds., Climate Change 2001: The Scientific Basis (Cambridge Univ.
657	Press, New York 2001).
658	Hynes, R. G.; Mössinger, J.; Cox, R. A. The interaction of HCl with water-ice at tropospheric
659	temperatures, Geophys. Res. Lett. 2001, 28, 2827-2830.
660	Iannarelli, R.; Rossi, M. J. $H_2O$ and HCl trace gas kinetics on crystalline and amorphous HCl
661	hydrates in the range 170 to 205 K: The HCl/H2O Phase Diagram revisited, Atm.
662	Chem. Phys. 2014, 14, 5183-5202.
663	Iannarelli, R.; Rossi, M. J. The mid-IR absorption cross sections of $\alpha\text{-}$ and $\beta\text{-}NAT$
664	(HNO <sub>3</sub> $\bullet$ 3H <sub>2</sub> O) in the range 170 to 185 K and of metastable NAD (HNO <sub>3</sub> $\bullet$ 2H <sub>2</sub> O) in
665	the range 172-182 K, J. Geophys. Res. Atmos. 2016, 120, 11707-11727.
666	Iannarelli, R.; Rossi, M. J. Heterogeneous Kinetics of $H_2O$ , $HNO_3$ and $HCl$ on $HNO_3$ hydrates
667	(α-NAT, β-NAT, NAD) in the range 175-200 K, Atm. Chem. Phys. 2016a 16, 11937-
668	11960.
669	Jensen, E. J.; Toon, O. B.; Vay, S. A.; Ovarlez, J.; May, R.; Bui; T. P.; Twohy, C. H.;
670	Gandrud, B. W.; Pueschel, R. F., Schumann, U. Prevalence of ice-supersaturated
671	regions in the upper troposphere: Implications for optically thin ice cloud formation,
672	J. Geophys. Res. 2001, 106, 17253-17266.
673	Kang, H.; Shin, T. H.; Park, S. P.; Kim, I. K.; Han, S. J. Acidity of Hydrogen Chloride on Ice
674	J. Am. Chem. Soc. 2000, 122, 9842-9843.
675	Leu, MT.; Moore, S.B.; Keyser, L.F. Heterogeneous Reactions of Chlorine Nitrate and
676	Hydrogen Chloride on Type I Polar Stratospheric Clouds J. Phys. Chem. 1991, 95,
677	7763-7771.
678	Lewellen, D.C. Persistent Contrails and Contrail Cirrus. Part II: Full Lifetime Behavior, J.
679	Atmos. Sci. 2014, 71, 4420-4438.
680	Lu, Q. B.; Sanche, L. Large enhancement in dissociative electron attachment to HCl
681	adsorbed on ice via transfer of presolvated electrons J. Chem. Phys. 2001, 115, 5711-
682	5713.
683	Lundgren, J. O.; Olovson, I. Hydrogen Bond Studies. XV. The Crystal Structure of Hydrogen
684	Chloride Dihydrate, Acta Cryst. <b>1967</b> , <i>23</i> , 966-970.
685	Lundgren, J. O.; Olovson, I. Hydrogen Bond Studies. XVI. The Crystal Structure of
686	Hydrogen Chloride Trihydrate, Acta Cryst. <b>1967a</b> , 23, 971-976.
687	Marcy, T. P; Fahey, D. W.; Gao, R. S.; Popp, P. J.; Richard, E. C.; Thompson, T. L.;
688	Rosenlof, K. H.; Ray, E. A.; Salawitch, R. J.; Atherton, C. S.; Bergmann, D. J.;
689	Ridley, B. A.; Weinheimer, A. J.; Loewenstein, M.; Weinstock, E. M.; Mahoney, M. J.





690	Quantifying Stratospheric Ozone in the Upper Troposphere with in situ Measurements
691	of HCl, Science <b>2004</b> , <i>304</i> , 261-265.
692	Marti, J.; Mauersberger, K. A survey and new measurements of ice vapor pressure at
693	temperatures between 170 and 250 K, Geophys. Res. Lett. 1993, 20, 363-366.
694	Mauersberger, K.; Krankowsky, D. Vapor pressure above ice at temperatures below 170 K,
695	Geophys. Res. Lett. 2003, 30, 1121-1124.
696	Oppliger, R.; Allanic, A., Rossi, M. J. Real-Time Kinetics of the Uptake of HOBr and
697	BrONO <sub>2</sub> on Ice and in the Presence of HCl in the Temperature Range 190 – 200 K, J.
698	Phys. Chem. A <b>1997</b> , 101, 1903-1911.
699	Ortega, I. K.; Escribano, R.; Fernandez-Torre, D.; Herrero, V. J.; Maté, B.; Moreno, M. A.
700	The HCl hexahydrate: RAIR spectra and theoretical investigation, Chem. Phys. Lett.
701	<b>2004</b> , <i>396</i> , 335-340.
702	Parent, P.; Laffon, C. Adsorption of HCl on the Water Ice Surface Studied by X-ray
703	Absorption Spectroscopy, J. Phys. Chem. B 2005, 109, 1547-1553.
704	Petrenko, V. F.; Whitworth, R. W. The Physics of Ice, Oxford University Press, 1999.
705	Pratte, P.; van den Bergh, H.; Rossi, M. J. The kinetics of H <sub>2</sub> O vapor condensation and
706	evaporation on different types of ice in the range 130-210 K, J. Phys. Chem. A 2006,
707	110, 3042-3058.
708	Schumann, U.; Baumann, R.; Baumgardner, D.; Bedka, S.T.; Duda, D.P.; Freudenthaler, V.;
709	Gayet, JF.; Heymsfield, A.J.; Minnis, P.; Quante, M.; Raschke, E.; Schlager, H.;
710	Vazquez-Navarro, M.; Voigt, C.; Wang, Z. Properties of individual contrails: A
711	compilation of observations and some comparisons, Atm. Chem. Phys. 2017, 17, 403-
712	438.
713	Schumann, U.; Kiemle, C.; Schlager, H.; Weigel, R.; Borrmann, S.; D'Amato, F.; Krämer,
714	M.; Matthey, R.; Protat, A.; Voigt, C.; Volk, M. Long-lived contrails and convective
715	cirrus above the tropical tropopause, Atm. Chem. Phys. 2017a, 17, 2311-2346.
716	Seinfeld, J. H.; Pandis, S. N. Atmospheric Chemistry and Physics, from Air Pollution to
717	Climate Change, John Wiley and Sons, Inc. (1998).
718	Solomon, S.; Garcia, R. R.; Rowland, F. S.; Wuebbles, D. J. On the Depletion of Antarctic
719	Ozone, Nature 1986, 321, 755-758.
720	Solomon, S.; Borrmann, S.; Garcia, R. R.; Portmann, R.; Thomason, L.; Poole, L. R.; Winker,
721	D.; McCormick, M. P. Heterogeneous chlorine chemistry in the tropopause region, J.
722	Geophys. Res. (Atmos) 1997, 102, 21411-21429.





723	Taesler, I.; Lundgren, J. O. Hydrogen Bond Studies. CXXIX. An X-Ray Determination of the
724	Crystal Structure of Hydrogen Chloride Hexahydrate, H <sub>9</sub> O <sub>4</sub> <sup>+</sup> Cl <sup>-</sup> •2H <sub>2</sub> O, Acta
725	Crystallogr. B 1978, 34, 2424-2428.
726	Tolbert, M. A.; Rossi, M. J.; Malhotra, R.; Golden, D. M. Reaction of Chlorine Nitrate with
727	Hydrogen Chloride and Water at Antarctic Stratospheric Temperatures, Science 1987,
728	238, 1258-1260.
729	Uras, N.; Rahman, M.; Devlin, J. P. Covalent HCl at the Surface of Crystalline Ice at 125 K:
730	The Stable Phase at Submonolayer Levels J. Phys. Chem. B 1998, 102, 9375-9377.
731	WMO (World Meteorological Organization) Scientific Assessment of Ozone Depletion 2002,
732	Global Ozone Research and Monitoring Project, report no. 47, Geneva, 2003.
733	Xueref, I.; Dominé, F. FTIR spectroscopic studies of the simultaneous condensation of HCl
734	and H <sub>2</sub> O at 190K – Atmospheric applications Atmos. Chem. Phys. 2003, 3, 1779-
735	1789.
736	Yoon, Y. K.; Carpenter, G. B. The Crystal Structure of Hydrogen Chloride Monohydrate,
737	Acta Cryst. 1959, 12, 17-20.
738	Zerefos, C. S.; Eleftheratos, K.; Balis, D. S.; Zanis, P.; Tselioudis, G.; Meleti, C. Evidence of
739	impact of aviation on Cirrus cloud formation, Atmos. Chem. Phys. 2003, 3, 1633-
740	1644.
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Table 1: Hardware parameters of both cryogenic sample supports	Si Optical Window	QCM	
Reactor temperature Tr [K]	32(	0	
Reactor volume V <sub>r</sub> [cm <sup>-3</sup> ]	235	0	
Conversion factor (1/RT) Conv [molec cm <sup>-3</sup> Torr <sup>-1</sup> ] with $R=62398$ [Torr cm <sup>3</sup> mol <sup>-1</sup> K <sup>-1</sup> ]	3.0 • 10	0 <sup>16 (1)</sup>	
Sample surface area [cm <sup>2</sup> ]	0.78	0.50	
$H_2O$ collision frequency with ice sample $\omega_{H_2O}$ [s <sup>-1</sup> ]	5.08	3.26	
$H_2O$ effusion rate constant of calibrated leak $k_{esc}(H_2O)$ [s <sup>-1</sup> ]	0.06	54	
MS calibration factor for $H_20$ (m/z=18, Stirred Flow) $C_{18}^{2}P^{10W}$ [molec s <sup>-1</sup> A <sup>-1</sup> ]	2.4 • 1	$10^{24}$	
MS calibration factor for $H_20$ (m/z=18, Dynamic ) $C_{10}^{47n}$ [molec s <sup>-1</sup> A <sup>-1</sup> ]	1.7 • 1	10 <sup>25</sup>	
HCl collision frequency with ice sample $\omega_{\rm HCl}$ [s <sup>-1</sup> ]	3.59	2.31	
HCl effusion rate constant of calibrated leak kese(HCl) [s <sup>-1</sup> ]	0.04	17	
MS calibration factor for HC1 (m/z=36, Stirred Flow) Cathered [molec s <sup>-1</sup> A <sup>-1</sup> ]	3.9 • 1	10 <sup>24</sup>	
MS calibration factor for HC1 (m/z=36, Dynamic) $C_{36}^{4ym}$ [molec s <sup>-1</sup> A <sup>-1</sup> ]	6.3 • 1	10 <sup>24</sup>	
Calculated escape orifice area A <sub>esc</sub> [mm <sup>2</sup> ]	1.0		
	$d = 10^4 \text{ Å}$		
	or 1.0 μm for O.D.= 1.08 [A] <sup>(2)</sup> at 3260 cm <sup>-1</sup>	Calibration Fe	actor
		Temperature [K]	ratio <sup>(3)</sup>
<sup>1</sup> Wall temperature of the reactor at $T = 320$ K		170	9.0
<sup>2</sup> See Reference [41]		180	8.0
<sup>3</sup> Corresponds to the ratio between the true number of molecules present on the QCM		190	7.8
support and the number of molecules displayed by the IC5 controller [42]		193	6.0
		205	2.0
		208	61





				experimen	ut and	l identifie	s the (	data displ	layed in Fi	igure 2			
Experiment	T <sub>ice</sub>	$d_0$	N <sup>0</sup> <sub>H20</sub>	R <sub>HCI</sub>	$\mathbf{t}_{\mathrm{dep}}$	N <sup>dep</sup>	HCI	N evap	$\chi^{0}_{HCl}$	$d_{\mathrm{D}}$	$\int_{eV}^{b}$	J <sup>e</sup>	r <sup>b/e</sup>
number	[K]	[Å]	[molec]	[molec s <sup>-1</sup> ]	[s]	[molec]	ML	[molec]		[Å]	[molec cm <sup>-2</sup> s <sup>-1</sup> ]	[molec cm <sup>-2</sup> s <sup>-1</sup> ]	
10	174	15230	$2.4 \cdot 10^{18}$	$6.4 \cdot 10^{12}$	94	$6.0 \cdot 10^{14}$	4.8	$4.7 \cdot 10^{14}$	$2.5 \cdot 10^{-4}$	2733	$1.9 \cdot 10^{15}$	$4.4 \cdot 10^{14}$	4.3
S.	188	13318	$2.0 \cdot 10^{18}$	$1.3 \cdot 10^{13}$	99	$8.7 \cdot 10^{14}$	7.0	8.9 · 10 <sup>14</sup>	$4.4 \cdot 10^{-4}$	4540	$1.2 \cdot 10^{16}$	3.9 · 10 <sup>15</sup>	3.1
4	190	14016	$2.1 \cdot 10^{18}$	$4.2 \cdot 10^{13}$	126	$5.4 \cdot 10^{15}$	43.2	3.6 · 10 <sup>15</sup>	$2.6 \cdot 10^{-3}$	6360	2.9 · 10 <sup>16</sup>	$1.4 \cdot 10^{15}$	20.7
9	190	13886	$2.1 \cdot 10^{18}$	$3.9 \cdot 10^{13}$	56	$2.2 \cdot 10^{15}$	17.6	1.8 • 10 <sup>15</sup>	$1.0 \cdot 10^{-3}$	12861	$3.4 \cdot 10^{16}$	$1.7 \cdot 10^{16}$	2.0
1	192	14926	$2.3 \cdot 10^{18}$	$3.1 \cdot 10^{12}$	36	$1.0 \cdot 10^{14}$	0.8	$1.8 \cdot 10^{14}$	$4.3 \cdot 10^{-5}$	2823	$2.9 \cdot 10^{16}$	$7.1 \cdot 10^{14}$	40.8
7	192	14682	$2.3 \cdot 10^{18}$	$8.0 \cdot 10^{11}$	356	$2.6 \cdot 10^{14}$	2.1	$1.6 \cdot 10^{14}$	$1.1 \cdot 10^{-4}$	6817	$3.2 \cdot 10^{16}$	$6.5 \cdot 10^{14}$	49.2
11	192	14420	$2.2\cdot 10^{18}$	$5.4 \cdot 10^{12}$	108	$5.4 \cdot 10^{14}$	4.3	$6.8 \cdot 10^{14}$	$2.4 \cdot 10^{-4}$	7717	$4.0 \cdot 10^{16}$	$7.9 \cdot 10^{14}$	50.6
3	193	14423	$2.2\cdot 10^{18}$	$3.5 \cdot 10^{12}$	220	$7.0 \cdot 10^{14}$	5.6	$8.1 \cdot 10^{14}$	$3.2 \cdot 10^{-4}$	5659	$4.9 \cdot 10^{16}$	$1.8 \cdot 10^{15}$	27.2
7	195	12614	$1.9 \cdot 10^{18}$	$4.3 \cdot 10^{12}$	45	$1.9 \cdot 10^{14}$	1.5	$1.8 \cdot 10^{14}$	$1.0 \cdot 10^{-4}$	5325	$4.6 \cdot 10^{16}$	$2.0 \cdot 10^{15}$	23.0
8	205	13505	$2.1 \cdot 10^{18}$	$1.6 \cdot 10^{13}$	36	$5.9 \cdot 10^{14}$	4.7	$3.0 \cdot 10^{14}$	$2.8 \cdot 10^{-4}$	4607	2.0 · 10 <sup>17</sup>	$1.0 \cdot 10^{16}$	20.0
6	210	13134	$2.0 \cdot 10^{18}$	$3.5 \cdot 10^{12}$	84	$3.0 \cdot 10^{14}$	2.4	$1.9 \cdot 10^{14}$	$1.5 \cdot 10^{-4}$	12136	$3.0 \cdot 10^{17}$	$1.8 \cdot 10^{16}$	16.7

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Table 2 : Representative experimental results for the kinetics of  $H_2O$  evaporation in the presence of HCl for increasing HCl deposition temperatures at given rates of deposition  $R_{HCl}$  and doses of HCl  $N_{HCl}^{dep}$ . In the first column the number refers to the corresponding





#### **Figure Captions**

Figure 1: Typical experimental protocol of the evaporation at 192 K of an approximately 1.2  $\mu$ m thick ice film doped with 5.4·10<sup>14</sup> molecules of HCl. This illustration corresponds to experiment 11 of Table 2. ( $\circ$ ): ice thickness monitored by QCM (Å), ( $\Box$ ): "apparent" H<sub>2</sub>O evaporative flux, J<sup>QCM</sup><sub>ev</sub>, monitored by QCM (molecule cm<sup>-2</sup> s<sup>-1</sup>), (+): I<sup>18</sup> MS signal for H<sub>2</sub>O, (×): I<sup>36</sup> MS signal for HCl (A), ( $\diamond$ ): J<sup>18</sup><sub>ev</sub> evaporative flux calculated from I<sup>18</sup> (molecule cm<sup>-2</sup> s<sup>-1</sup>), ( $\Delta$ ): Int(J<sup>18</sup><sub>ev</sub>) time integral of J<sup>18</sup><sub>ev</sub> (molecule cm<sup>-2</sup>).

Figure 2: Change of the evaporative flux  $J_{ev}(H_2O)$  as a function of the HCl mole fraction ( $\chi_{HCl}$ ) for the cases presented in Table 2 color-coded according to the corresponding experiment number in Table 2. The colored and circled numbers on axis "b" (left) correspond to  $J_{ev}(H_2O)$  of pure ice before HCl deposition, the ones on axis "e" (right) are  $J_{ev}(H_2O)$  at  $t = t_{He}$  at the end of HCl evaporation. The colored circles in the data field mark the value of  $J_{ev}(H_2O)$  after HCl deposition at  $t = t_0$  and are equal to  $J_{ev}(H_2O)$  of pure ice. The start of any particular  $J_{ev}(H_2O)$  curve as a continuous solid (bold) line occurs at  $t = t_D$  at 85% of  $J_{ev}(H_2O)$  at t = 0 (pure ice value, colored circle or circled number on axis "b" to the left) and ends at  $t_{Hb}$ , the beginning of HCl evaporation as displayed in Figure 1B.

Figure 3: Synopsis of the dependence of the evaporation range parameter  $r^{b/e}$  on the rate of deposition  $R_{HCl}$  of HCl for temperatures between 188 and 210 K. Each point is marked with the total number of HCl molecules ( $N_{HCl}$ ) deposited on the ice film , the temperature of the ice film at HCl deposition and the experiment number (bold) referring to Table 2. The hashed area encompasses  $r^{b/e}$  values for dataset B (experiments 3,4,7,8). The color code goes from low (blue) over medium (green) to high (red) temperatures.

Figure 4: Synopsis of the dependence of  $d_D$  on the rate of deposition  $R_{HCl}$  of HCl for temperatures between 188 and 210 K. Each point is marked with the total number of HCl molecules ( $N_{HCl}$ ) deposited on the ice film, the temperature of the ice film at HCl deposition and the experiment number (bold font) referring to Table 2. The hashed area encompasses  $d_D$  values for dataset B (experiments 3,4,7,8). The color code goes from low (blue) over medium (green) to high (red) temperatures.







Time [s]





















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Figure A1: Synopsis of the dependence of the evaporation range parameter, r<sup>b/e</sup>, on the number of adsorbed HCI, N<sub>HCl</sub><sup>dep</sup>, adsorbed on ice for temperatures between 188 and 210 K. each point is marked with the deposition rate of HCl molecules in molec s<sup>-1</sup> on the ice film, the temperature of the ice film and the experiment running number (bold) referring to Table 2.

APPENDIX A: Figures A1, A2, A3 and A4







Figure A2: Synopsis of the dependence of the remaining thickness d<sub>D</sub> on the number of adsorbed HCl, N<sub>HCl</sub><sup>dep</sup>, dispensed on ice for temperatures between 188 and 210 K. Each point is marked with the deposition rate of HCl in molec s<sup>-1</sup> on the ice film, the temperature of the ice film and the experiment running number (bold) referring to Table 2.







Figure A3: Synopsis of kinetic and thermodynamic results for an amorphous  $H_2O/HCl$  mixture using HCl as a probe gas. The symbols/colors used correspond to different experimental runs and the graphs show the scatter of the individual measurements within a series. Original data are published in Iannarelli, R. and Rossi, M.J. Atmos. Chem. Phys. 14, 5183–5204, 2014.







Figure A4: Synopsis of kinetic and thermodynamic results for crystalline HCl hexahydrate (HH) using X = HCl as a probe gas. The symbols/colors used correspond to different experimental runs and the graphs show the scatter of the individual measurements within a series. Original data are published in Iannarelli, R. and Rossi, M.J. Atmos. Chem. Phys. 14, 5183–5204, 2014.