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2 **The Influence of HCl on the Evaporation Rates of H₂O over Water Ice in**
3 **the Range 188 to 210 K at small Average Concentrations**

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19 **ABSTRACT**

20
21 The evaporation flux $J_{\text{ev}}(\text{H}_2\text{O})$ of H₂O from HCl-doped typically 1.5 μm or so thick vapor-
22 deposited ice films has been measured in a combined quartz crystal microbalance (QCMB) – residual
23 gas mass spectrometry (MS) experiment. $J_{\text{ev}}(\text{H}_2\text{O})$ has been found to show complex behaviour and to
24 be a function of the average mole fraction χ_{HCl} of HCl in the ice film ranging from 6×10^{14} to 3×10^{17}
25 molecule cm⁻² s⁻¹ at 174 – 210 K for initial values χ_{HCl}^0 ranging from 5×10^{-5} to 3×10^{-3} at the start of the
26 evaporation. The dose of HCl on ice was in the range of 1 to 40 formal monolayers and the H₂O vapor
27 pressure was independent of χ_{HCl} within the measured range and equal to that of pure ice down to 80
28 nm thickness. The dependence of $J_{\text{ev}}(\text{H}_2\text{O})$ with increasing average χ_{HCl} was correlated with (a) the
29 evaporation range $r^{\text{b/e}}$ parameter, that is the ratio of $J_{\text{ev}}(\text{H}_2\text{O})$ just before HCl-doping of the pure ice
30 film and $J_{\text{ev}}(\text{H}_2\text{O})$ after observable HCl desorption towards the end of film evaporation, and (b) the
31 remaining thickness d_{D} below which $J_{\text{ev}}(\text{H}_2\text{O})$ decreases to less than 85% of pure ice. The dependence
32 of $J_{\text{ev}}(\text{H}_2\text{O})$ with increasing average χ_{HCl} from HCl-doped ice films suggests two limiting data sets, one
33 associated with the occurrence of a two-phase pure ice/crystalline HCl hydrate binary phase (set A),
34 and the other with a single phase amorphous HCl/H₂O binary mixture (set B). The measured values of
35 $J_{\text{ev}}(\text{H}_2\text{O})$ may lead to significant evaporative life-time extensions of HCl-contaminated ice cloud
36 particles under atmospheric conditions, regardless of whether the structure corresponds to an
37 amorphous or crystalline state of the HCl/H₂O aggregate.

40 1. INTRODUCTION

41

42 HCl is among the mineral acids that control the acidity of the atmosphere together
43 with HNO₃ and H₂SO₄. The production of atmospheric HCl is predominantly taking place in
44 the middle and upper stratosphere where O₃ is formed owing to photolysis of halogen
45 containing source gases such as CFC's (chlorofluorocarbons). However, there are no known
46 sources of HCl in the upper troposphere (UT) because scavenging processes of HCl
47 throughout the troposphere are very efficient which leads to HCl background concentrations
48 of less than 0.1 ppb (Graedel and Keene, 1995). The absence of significant sources in the
49 troposphere, the long photolytic lifetime of HCl and the fact that the production region is well
50 separated from the regions of interest, namely the UT and the lower stratosphere (LS) all
51 contribute to the fact that HCl is an excellent tracer for stratospheric ozone in the UT (Marcy
52 et al., 2004). Owing to the frequent occurrence of Cirrus clouds in this atmospheric region it
53 is of obvious interest to study the interaction of HCl with atmospheric ice particles at relevant
54 temperature and pressure conditions (Jensen et al. 2001; Zerefos et al., 2003). The compact
55 correlation between O₃ and HCl has been used to monitor stratospheric-tropospheric
56 exchange processes and stratospheric O₃ intrusions into the troposphere that are still an active
57 field of investigation (Houghton et al, 2001).

58 HCl is of importance in the LS as it partakes in heterogeneous reactions on Polar
59 Stratospheric Ice Clouds (PSC's) as well as on background stratospheric H₂SO₄ aerosol
60 according to the following reaction taken as an example:



62 These reactions efficiently convert inactive Cl-containing reservoir molecules such as HCl
63 and ClONO₂ into active photolyzable Cl-containing compounds in a single reaction. Typical
64 examples of such photolabile reaction products are Cl₂, ClNO₂ and HOCl that will change the
65 atmospheric composition owing to the high reactivity of the photolysis products such as
66 atomic Cl (Solomon et al., 1986; Tolbert et al., 1987; WMO 2003). It thus follows that HCl is
67 of stratospheric importance and is frequently used as a model compound for heterogeneous
68 reactions on ices that has inspired many laboratory kinetic studies (Leu et al., 1991; Hanson
69 and Ravishankara, 1992; Chu et al., 1993; Flückiger et al., 1998; Hynes et al., 2001; Abbatt,
70 2003).

71 HCl forms hydrates of variable stoichiometry when exposed to ice depending on the
72 temperature of deposition and the partial pressure of HCl (Ortega et al., 2004; Graham and
73 Roberts, 1997). X-Ray diffraction has allowed the identification of four crystalline hydrates

74 containing one (Yoon and Carpenter, 1959), two (Lundgren and Olovson, 1967), three
75 (Lundgren and Olovson, 1967a) and six (Taesler and Lundgren, 1978) H₂O per HCl molecule.
76 In addition, amorphous mono-, tetra- and hexahydrates have been reported under various
77 experimental conditions (Yoon and Carpenter, 1959; Delzeit et al., 1993). The control of
78 growth conditions of a specific HCl hydrate is sometimes elusive, but the formation of a
79 saturated HCl hexahydrate phase has been reported at sufficiently large HCl exposure
80 (Graham and Roberts, 1995) using amorphous ice as a starting point despite the fact that the
81 hexahydrate is said to nucleate with difficulty, at least in thin films (Ortega et al., 2004).
82 However, the molecular and dynamic details of the crystallization process have not been
83 investigated as yet.

84 Fourier-Transform IR (FTIR) absorption measurements have enabled the
85 characterization of both amorphous as well as crystalline HCl hydrates at growth conditions
86 that are sometimes significantly different compared to the samples investigated using X-Ray
87 diffraction. Vibrational spectra of HCl hydrates in the mid IR have been routinely used for
88 identification purposes for some time (Ferriso and Hornig, 1955; Gilbert and Sheppard,
89 1973). Recently, the mid IR absorption spectra of the four HCl hydrates mentioned above
90 have been assigned in a comprehensive and definitive way, albeit without simultaneous proof
91 of the crystalline structure using X-Ray diffraction (Buch et al., 2002; Xueref and Dominé,
92 2003). More recently, the reflection absorption IR spectrum (RAIR) of crystalline HCl
93 hexahydrate in the mid-IR range has been recorded and assigned using theoretical calculations
94 based on density functional theory that results in a refinement of the geometric structure of
95 the HCl hydrates and a prediction of the vibrational modes of the crystal (Ortega et al., 2004).
96 It must be recalled that FTIR spectra in transmission and reflection may in most cases not be
97 directly compared across the mid IR range.

98 Regarding the nature of the HCl-ice adsorbate one of the important questions is
99 whether adsorbed HCl is ionized or exists as a molecular adsorbate under atmospherically
100 relevant conditions of the UT/LS. This will determine the mechanism of the heterogeneous
101 reaction which constitutes necessary knowledge for the extrapolation of heterogeneous
102 reaction rates measured in the laboratory to atmospheric conditions. Thermal desorption of
103 HCl monitored by IR absorption in the mid-IR range revealed a molecularly adsorbed state of
104 HCl desorbing below 50 K (Delzeit et al., 1993a). IR studies performed by Banham et al. on
105 HCl-ice films failed to detect molecularly adsorbed HCl at $T \geq 90$ K despite the high rate of
106 HCl adsorption in that temperature range (Banham et al., 1995). In contrast, Graham and
107 Roberts attributed a characteristic Temperature Programmed Desorption (TPD) peak of a

108 HCl/amorphous ice adsorbate monitored by residual gas MS and occurring at 150 K to
109 molecularly adsorbed HCl (Graham and Roberts, 1995). However, they did not report the IR
110 absorption spectrum of the adsorbate in the mid-IR nor did they explain why molecular
111 adsorption of HCl exclusively occurred on amorphous, but not on crystalline ice. Most recent
112 results seem to point towards the existence of molecularly adsorbed HCl on ice below 50 K
113 and at submonolayer coverages in coexistence with ionized solvated HCl whose fraction
114 increases with increasing ice temperature (Buch et al., 2002; Delzeit et al., 1993a; Delzeit et al.
115 1997; Uras et al., 1998; Devlin et al., 2002; Lu and Sanche, 2001). Kang et al. discovered
116 that both molecularly adsorbed as well as ionized HCl coexisted on ice that was deposited
117 under Ultra-High Vacuum (UHV) conditions in the temperature range 50 to 140 K and under
118 conditions of low HCl exposure (Kang et al., 2000).

119 Although theoretical electronic structure calculations predict spontaneous ionization of
120 adsorbed HCl (Gertner and Hynes, 1996; Bolton and Petterson, 2001) most experiments point
121 towards a seemingly thermally activated ionization process that may be enabled by structural
122 factors of the ice matrix that are themselves a function of temperature. Consistent with these
123 results concentration profiling experiments of HCl/ice adsorbates using static Secondary
124 Ionization Mass Spectroscopy (SIMS) techniques failed to discover molecularly adsorbed
125 HCl on ice in the range 90-150 K (Donsig and Vickerman, 1997). In conclusion, both
126 experimental and theoretical studies clearly point to the absence of significant quantities of
127 molecularly or covalently adsorbed HCl under stratospheric conditions. Instead, HCl is
128 ionized and solvated by H₂O on the surface of ice films and may occur either as amorphous
129 HCl/H₂O hydrates of undefined stoichiometry or as crystalline HCl hydrates. However, these
130 facts do not rule out the presence of small amounts of molecularly adsorbed HCl on ice that
131 may be intermediates in the complex mechanism of HCl adsorption on ice as evidenced by
132 the negative temperature dependence of the rate of uptake of HCl on ice (Flückiger et al.,
133 1998). In fact, such an intermediate has been invoked in the description of HCl adsorption on
134 ice under atmospheric conditions using a chemical kinetic model based on a multitude of
135 experimental observables collected upon HCl uptake on ice (Flückiger and Rossi, 2003).

136 Work by Parent and coworkers uses Near-Edge X-Ray Absorption Spectroscopy
137 (NEXAFS) of HCl-doped low temperature ice substrates in order to determine the relative
138 population of ionic and covalently bound HCl and distinguish between bulk and HCl surface
139 states in the temperature range 20 to 150 K (Bournel et al., 2002; Parent and Laffon, 2005).
140 The results seem to confirm the consensus on the low-temperature existence of molecularly
141 adsorbed HCl up to 90 K beyond which an increasing amount of HCl is converted into an

142 ionic form, such as $\text{H}_3\text{O}^+\text{Cl}^-$ (Eigen cation) or $\text{H}_5\text{O}_2^+\text{Cl}^-$ (Zundel cation) formed by
143 spontaneous ionization of adsorbed HCl on ice, up to completion at 150 K (). The newest
144 work by Parent compares NEXAFS with photoemission (UPS, XPS) and FTIR in
145 transmission of thin HCl/H₂O films (Parent et al., 2011). The results are roughly consistent
146 but surprising in the sense that these workers find 92% ionically dissolved HCl in/on ice at 50
147 K in contrast to Kang et al. (2000) and Devlin et al. (2000) under similar exposure (dose) and
148 temperature conditions. In addition, Parent et al. (2011) perform the NEXAFS experiment on
149 a (thick) 100 ML “crystalline” H₂O ice substrate deposited at 150 K whereas the
150 photoemission and FTIR absorption experiments used a 4 ML thin ice slab deposited at 120
151 K. The question has to be raised whether the two types of used ice films may be responsible
152 for some of the discrepancies in the results because both the density and the structure of ice
153 are known to be a strong function of temperature and deposition conditions (Kuhs et al., 2012;
154 Schriver-Mazzuoli et al., 2000). The most recent work of Parent et al. (2011) sparked an
155 interesting controversy in the assignment of the FTIR absorption spectrum of thin HCl/H₂O
156 films and led to two comments showcasing the difficulties of intercomparison of nominally
157 identical experiments (Devlin and Kang, 2012; Parent et al., 2012).

158 Furthermore, the results indicate that the “dangling bonds” of the ice surface attributed
159 to isolated OH groups are not the unique site of HCl adsorption, even in the range 20-90 K
160 (Flückiger and Delval, 2002). The present work suggests that maiden uptake of HCl onto
161 pure ice weakens and perturbs the crystal structure of the ice matrix in an irreversible way
162 such that additional sites for HCl adsorption and ionization are created akin to Parent et al.
163 (2011). Initial HCl uptake on pure ice therefore has a catalytic effect on the following HCl
164 uptake. This irreversible nature of initial HCl dosing is known for several years and has been
165 observed some time ago in Knudsen flow reactor studies on the HCl/H₂O system under
166 steady-state conditions of both HCl and H₂O at temperatures representative of the UT/LS
167 (Flückiger et al., 1998; Oppliger et al., 1997). The most recent experimental work on
168 HCl/H₂O at an atmospherically relevant (“warm”) temperature (253 K) has examined the HCl
169 depth profile using XPS spectroscopy and finds molecularly adsorbed (physisorbed) HCl at its
170 outermost layer and ionic dissociation in deeper layers (Kong et al., 2017). Complementary
171 X-Ray absorption results also point towards a perturbation of the crystal structure of ice in the
172 aftermath of HCl adsorption/dissolution into deeper layers of ice.

173 We have concluded from recent work that HCl doping in quantities of submonolayer
174 to several monolayers of HCl leads to the decrease of both the evaporative flux J_{ev} (molecule
175 $\text{cm}^{-2}\text{s}^{-1}$) or rate R_{ev} (molecule $\text{cm}^{-3}\text{s}^{-1}$) and the rate of condensation k_{cond} (s^{-1}), of H₂O in the

176 presence of ice without perturbing the equilibrium vapour pressure of H₂O, $P_{\text{H}_2\text{O}}^{\text{eq}}$ (Delval et
177 al., 2003). We have furthermore shown that the way J_{ev} of H₂O decreases with time depends
178 on the rate of deposition or the integral of deposited HCl, namely R_{HCl} (molecule s⁻¹) and N_{HCl}
179 (molecule), respectively. It appears that two observed HCl species on/in ice, namely single
180 phase amorphous HCl/H₂O mixtures and a binary phase consisting of pure ice and an as yet
181 unidentified crystalline HCl hydrate, HCl•xH₂O, decrease $J_{\text{ev}}(\text{H}_2\text{O})$ to a different extent as
182 proposed in Delval et al. (2003). These results have led us to perform systematic experiments
183 in this work using the Quartz Crystal MicroBalance (QCMB) combined with residual gas
184 Mass Spectrometry (MS) that we have used successfully in the past (Delval and Rossi, 2004)
185 in order to investigate the temporal change of $J_{\text{ev}}(\text{H}_2\text{O})$ with the increasing average mole
186 fraction of HCl, χ_{HCl} , remaining in the ice. One of the goals of the present work is to
187 determine the influence of the HCl deposition parameters on the temporal change of J_{ev} and
188 the mass accommodation coefficient α during evaporation of a HCl-doped ice film and its
189 consequence on the lifetime of atmospheric ice particles contaminated by HCl. This issue is
190 key in relation to the importance of heterogeneous vs. homogeneous atmospheric reactions at
191 midlatitudes as has been pointed out in the past (Solomon et al., 1986; 1997).

192

193 **2. EXPERIMENTAL**

194

195 The emphasis of the present experiments was placed on the deposition of small
196 amounts of HCl ranging in doses from 1 to 40 formal monolayers of HCl where a formal
197 monolayer of adsorbed HCl corresponded to a surface concentration of 2.5×10^{14} molecule cm⁻²
198 (Table A1) which is a consensus value obtained from several selected experiments. The
199 apparatus as well as the methods used for calibration and the HCl deposition procedure have
200 been described in detail elsewhere (Delval and Rossi, 2005). The experimental conditions are
201 generally identical to the ones presented in Delval and Rossi (2005) and the instrumental
202 parameters are summarized in Table 1. The only significant difference between the study of
203 HNO₃-doped ice and the present condensed phase investigation of HCl-doped ice lies in the
204 mode of trace gas admission. HCl was deposited by backfilling the reactor under stirred flow
205 conditions with the inlet tubing used for trace gas injection oriented towards one side of the
206 Si-window of the cryostat set at ambient temperature whereas HNO₃ was deposited by
207 directed injection onto ice films supported by the quartz crystal of the QCMB as referenced
208 above. Evaporation experiments have been performed isothermally on samples in the

209 temperature range 174-210 K under dynamic pumping conditions, that is at maximum
210 pumping speed (gate valve open) in order to prevent readsorption of HCl on the ice substrate.

211 First, an approximately 1.5 μm thick ice film was grown at 190 K on the quartz crystal
212 of the QCM by deposition of bidistilled water vapor at a rate of 1×10^{17} molecule $\text{cm}^{-2} \text{s}^{-1}$
213 under static conditions. The H_2O equilibrium vapor pressure agreed with published values
214 across the covered temperature range (Marti and Mauersberger, 1993; Mauersberger and
215 Krankowsky, 2003). Subsequently, the system was set to the desired temperature given in
216 Table 2 (second column from the left) and a metered amount of HCl was deposited under
217 stirred flow conditions. The rate of deposition of HCl, R_{HCl} , as well as its time integral,
218 namely the number of HCl molecules deposited on ice, N_{HCl} , have been evaluated using the
219 method described in Delval and Rossi (2005). Typically, R_{HCl} ranges between 8.0×10^{11} and
220 4.2×10^{13} molecule s^{-1} and N_{HCl} between 1.0×10^{14} and 5.4×10^{15} molecules. The experimental
221 conditions of HCl-deposition as well as important experimental parameters are reported in
222 Table 2. Finally, the system was set to dynamic pumping conditions by opening the gate valve
223 to the turbopump. $J_{\text{ev}}(\text{H}_2\text{O})$ was measured isothermally using both the QCBM and residual gas
224 MS. Figure 1 illustrates a typical experimental protocol of the evaporation at 192 K of a HCl-
225 doped ice film labelled as experiment 11 in Table 2 and performed as a multidagnostic
226 experiment where both the gas- as well as the condensed phases are simultaneously
227 monitored.

228 At $t = 0$, the system is set from stirred flow to dynamic pumping that starts the
229 evaporation experiment. The continuous curve marked with the empty squares symbol in
230 Figure 1A corresponds to $J_{\text{ev}}^{\text{QCM}}$, the evaporative flux of H_2O calculated from the raw signal
231 at the output of the QCM. The diamond symbol (\diamond) corresponds to J_{ev}^{18} evaluated from I^{18} ,
232 the MS signal amplitude for H_2O monitored at $m/e = 18$. $\text{Int}(J_{\text{ev}}^{18})$ marked by triangles in
233 Figure 1A is the time integral of J_{ev}^{18} and corresponds to the total number of H_2O molecules
234 that have evaporated from the ice film at t . D is the label at time t_D at which $J_{\text{ev}}(\text{H}_2\text{O})$
235 decreased from its original value corresponding to pure ice to 85% of its original value at $t =$
236 0 , and d_D is the remaining thickness of the ice film at t_D . H_b and H_e in Figure 1B correspond to
237 the time when HCl evaporation begins and ceases to be observed, respectively, using gas
238 phase residual mass spectrometry (x symbols in Figure 1B) and are labeled t_{H_b} and t_{H_e} . The
239 data have been treated in analogy to HNO_3 -doped ice through the formalism given in Delval
240 and Rossi (2005). Akin to HNO_3 the mass balance between HCl deposited, $N_{\text{HCl}}^{\text{dep}}$, and HCl
241 recovered during ice evaporation, $N_{\text{HCl}}^{\text{evap}}$, agrees to within less than a factor of 2 under
242 dynamic pumping conditions. We therefore estimate the average uncertainty (2σ) of the HCl

243 mole fraction χ_{HCl} of $\pm 18\%$ from the average discrepancy between $N_{\text{HCl}}^{\text{dep}}$ and $N_{\text{HCl}}^{\text{evap}}$ displayed
244 in Table 2. In the following N_{HCl} will always refer to $N_{\text{HCl}}^{\text{dep}}$ derived from the measurement of
245 HCl at deposition because it refers to a directly measured quantity originating from a
246 measured pressure decrease in a given volume and time interval $\Delta P/\Delta t$. The present
247 experiments cover the evaporation of a small albeit important fraction of the model ice film
248 for which the decrease of $J_{\text{ev}}(\text{H}_2\text{O})$ is significant.

249

250 3. RESULTS

251

252 The experimental data reported in Table 2 on the isothermal change of the evaporative
253 flux of water, $J_{\text{ev}}(\text{H}_2\text{O})$, as a function of the average mole fraction of HCl, χ_{HCl} , in the
254 remaining ice film during the evaporation process under dynamic conditions are presented in
255 Figure 2. Dynamic pumping conditions ensure the absence of any readsorption of H_2O vapor
256 during evaporation owing to the low H_2O partial pressures in the reactor. The axes labelled
257 "b" and "e" correspond to the values of $J_{\text{ev}}(\text{H}_2\text{O})$ at the end of ice film deposition and after
258 desorption of most of the adsorbed HCl from the HCl-doped ice film at t_{He} , respectively, as
259 displayed in Figure 1B. The average mole fraction χ_{HCl} of HCl in the remaining ice film as a
260 function of time is calculated according to Delval and Rossi (2005). The change in χ_{HCl} owing
261 to H_2O evaporation is evaluated between $t = t_{\text{D}}$ and $t = t_{\text{Hb}}$ that corresponds to the time interval
262 when the number of adsorbed HCl molecules is constant as no release of HCl is observable in
263 the gas phase at $m/e = 36$ before t_{Hb} . Table 2 also displays the initial value of the HCl mole
264 fraction, χ_{HCl}^0 , calculated for the ice film just at the end of HCl deposition and marked by a
265 colored circle on the experimental trajectory of a color-coded evaporating ice film displayed
266 in Figure 2. The average mole fraction of HCl in the ice film, χ_{HCl} , increases owing to
267 evaporation of H_2O from the ice film without loss of HCl such that the elapsed time increases
268 with χ_{HCl} in Figure 2.

269 The beginning of an evaporation experiment after the end of HCl doping ($t = 0$ in
270 Figure 1 or t_0 in Figure 2) is marked by a colored circle of a given experiment whose
271 parameters are displayed in Table 2 and Figure 2 (see experiment 8). As pointed out above, at
272 $t = t_{\text{D}}$ $J_{\text{ev}}(\text{H}_2\text{O})$ has decreased to an arbitrarily chosen value of 85% of its original value
273 measured at $t = t_0$ that corresponds to the beginning of the bold color-coded smooth curve of a
274 given experiment. Figure 2 essentially displays trajectories of evaporation experiments from t_0
275 (colored circle) moving to t_{D} and finishing at t_{Hb} between the two limiting values for pure ice

276 (color coded number of a given experiment on axis “b” for “beginning”) and the remaining
277 ice film at the end of measurable HCl desorption t_{He} (color-coded number of experiment on
278 axis “e” for “Halogen end”). The trajectory of an experiment with values of χ_{HCl} between t_0
279 (colored circle at χ_{HCl}^0) and t_D (beginning of bold colored line, see experiment 8 in Figure 2)
280 ending at t_{Hb} (end of bold line, experiment 8) is presented as a bold dashed-dotted and bold
281 smooth line from t_0 to t_{Hb} , respectively, in order to emphasize the quantitative portion of the
282 experiment. Thinner (color-coded) dotted lines connect the end of ice film deposition (colored
283 circle on axis “b”) and HCl-dosing with t_0 , the beginning of the evaporation experiment and
284 also describe the post-phase of evaporation starting at t_{Hb} to t_{He} in order to guide the eye of the
285 reader to imagine a complete evaporation cycle.

286 Two different data sets of the change of $J_{ev}(H_2O)$ with χ_{HCl} may be distinguished in
287 Figure 2. The first kind of data set corresponds to the curves describing J_{ev} for experiments 1,
288 2, 9 and 11 and is called dataset A. These traces present a slow continuous decrease of
289 $J_{ev}(H_2O)$ as χ_{HCl} increases during H_2O evaporation. The second type of dataset shows an
290 initial plateau of $J_{ev}(H_2O)$ with increasing χ_{HCl} starting at the value of pure ice evaporation
291 followed by a sudden decrease of $J_{ev}(H_2O)$ and is found for experiments 3, 4, 7 and 8 which
292 we call dataset B. Akin to HNO_3 , we have evaluated the impact of the HCl deposition
293 protocol on the evaporation range parameter, $r^{b/e}$, which is the ratio between the evaporative
294 flux of H_2O at the beginning of ice evaporation, $J_{ev}^b(H_2O)$ reported on the left axis “b” in
295 Figure 2, and $J_{ev}(H_2O)$ close to the end of the desorption of HCl, $J_{ev}^e(H_2O)$, at $t = t_{He}$ (the right
296 axis “e” in Figure 2). It describes the factor by which $J_{ev}(H_2O)$ decreases within the limits of
297 “b” and “e”. The impact of both the rate of deposition of HCl on ice, R_{HCl} , and its time
298 integral corresponding to the dose of deposited HCl, N_{HCl} , are presented in Figures 3 and
299 Figure A1 (Appendix), respectively.

300 It appears from these Figures that we have not succeeded to find a simple experimental
301 parameter that controls $J_{ev}(H_2O)$ either with elapsed time or amount of adsorbed HCl
302 expressed as the time dependence of χ_{HCl} . Instead, the data may roughly be classified along
303 the two cases presented above, namely datasets A and B. The distinction between both data
304 sets seems to be the rate of change (slope) of $J_{ev}(H_2O)$ within a fairly narrow range of χ_{HCl} .
305 Indeed, the available number of experiments clearly shows two distinct and limiting cases
306 whereas the search for other controlling parameters such as R_{HCl} , N_{HCl} and the temperature of
307 deposition (T_{ice}) for dataset A failed akin to a similar HNO_3 study (Delval and Rossi, 2005).

308 One may take note for instance of the low value of χ_{HCl} at 210 K for experiment 9
309 where the conditions of deposition are similar to experiments 1 and 2, yet, its respective

310 values of $r^{b/e}$ differ significantly from experiment 9 (Figure 3). In contrast, for dataset B the
311 $r^{b/e}$ values are similar for the whole set and range from 20 to 27.2 staying within a fairly
312 narrow band. Moreover, they seem to be independent of R_{HCl} and N_{HCl} as for data set A. In
313 contrast, the $r^{b/e}$ values for dataset A seem widely scattered over the explored parameter space.
314 We have also investigated the impact of the deposition protocol on d_{D} , which is the thickness
315 of ice that is affected by the presence of HCl, namely the remaining thickness of ice whose
316 $J_{\text{ev}}(\text{H}_2\text{O})$ value has decreased to 85% of $J_{\text{ev}}(\text{H}_2\text{O})$ of pure ice. The results on d_{D} as a function
317 of R_{HCl} and $N^{\text{dep}}_{\text{HCl}}$ are presented in Figures 4 and A2 (Appendix), respectively. Taking the
318 results of Figures 3, 4, A1 and A2 together we arrive at the following two conclusions:
319 (1) T_{ice} , R_{HCl} and $N^{\text{dep}}_{\text{HCl}}$ are not controlling parameters or predictors for $J_{\text{ev}}(\text{H}_2\text{O})$ of either set.
320 (2) The evaporation range parameters $r^{b/e}$ and d_{D} are not characterizing set A. In contrast, for
321 dataset B, $r^{b/e}$ and d_{D} values fall into a narrow range with values varying from 460.7 to 636.0
322 nm compared to the original ice thickness d_0 of 1'500 nm or so (exact numbers in Table 2).

323

324 4. DISCUSSION

325

326 Figure 1 displays the evaporation history of sample 11 as an example whose deposition
327 parameters are listed in Table 1. The initial average mole fraction χ_{HCl}^0 of HCl, once
328 deposition on the 1.44 μm thick ice film under stirred flow reactor conditions is terminated,
329 has been estimated from the total number of H_2O molecules contained in the ice film and the
330 measured number of deposited HCl molecules, $N^{\text{dep}}_{\text{HCl}}$, for experiment 11 (Table 2). Table 2
331 and Figure 1 reveal that for approximately 2.2×10^{18} H_2O molecules in the film and 5.4×10^{14}
332 molecules of deposited HCl, we obtain $\chi_{\text{HCl}}^0 = 2.7 \times 10^{-4}$. This HCl mole fraction represents an
333 average value that takes into account all H_2O molecules contained in the ice film whereas in
334 reality there will be a HCl gradient across the ice film as has been observed in the case of the
335 HNO_3/ice system (Delval and Rossi, 2005).

336 After the HCl deposition process on the typically 1.5 μm thick ice film the gate valve is
337 opened in order to initiate the isothermal evaporation experiment under dynamic pumping
338 conditions. Initially, H_2O evaporates at fluxes $J_{\text{ev}}(\text{H}_2\text{O})$ that are characteristic of pure ice
339 measured previously (Delval and Rossi, 2004; Pratte et al., 2006). These initial values
340 $J_{\text{ev}}^b(\text{H}_2\text{O})$ are displayed on the left-hand “b” (= “beginning”) axis in Figure 2. As the
341 evaporation proceeds $J_{\text{ev}}(\text{H}_2\text{O})$ slightly decreases with time as displayed in Figure 1A to the
342 arbitrarily chosen point where $J_{\text{ev}}(\text{H}_2\text{O})$ has decreased to 85% of the initial pure ice value at
343 which point the remaining ice thickness d_{D} has decreased by approximately one third to 771.7

344 nm remaining ice thickness as displayed in Figure 1B and Table 2. Further evaporation of
345 H₂O leads to a continuous decrease of J_{ev}(H₂O) at a corresponding increase of χ_{HCl} up to point
346 H_b defined above (“Halogen beginning”) at t_{Hb} (Figure 1B) where HCl starts to desorb from
347 the ice film as monitored using the residual MS signal at m/e = 36.

348 For $t < t_{\text{Hb}}$, χ_{HCl} is given by the number of originally deposited HCl molecules that
349 remain adsorbed on the ice film up to t_{Hb} and the remaining H₂O molecules in the film. In
350 contrast, for $t > t_{\text{Hb}}$ the composition of the remaining ice film must be determined by taking
351 into account the loss by evaporation of both H₂O and HCl. The present experimental
352 configuration is not adapted to quantitatively measure HCl loss. Therefore, we have chosen to
353 display the temporal development of J_{ev}(H₂O) for $t < t_{\text{Hb}}$ in Figure 2 as a function of the
354 average value of the HCl mole fraction χ_{HCl} . However, the value of J_{ev}(H₂O) at $t = t_{\text{He}}$ where
355 most of the HCl has desorbed from the ice film is plotted on the right axis labelled “e” (=
356 “end”) as J_{ev}^e(H₂O) in Figure 2 in order to provide a limit for the minimum value of the
357 evaporation rate J_{ev}(H₂O) at an ice film thickness d_{He} of approximately 80 ± 10 nm as
358 displayed in Figure 1B. We have observed in the past that J_{ev}(H₂O) for a pure ice film of
359 approximate thickness of 80 nm or less also slows down, presumably owing to island
360 formation at the very end of pure thin ice film evaporation (Delval and Rossi, 2005).
361 Therefore, results are becoming more difficult to interpret such that we halted the experiment
362 at t_{He}. The ratio $r^{\text{b/e}} = J_{\text{ev}}^{\text{b}}(\text{H}_2\text{O})/J_{\text{ev}}^{\text{e}}(\text{H}_2\text{O})$ is displayed in Table 2 and is an operational
363 evaporation range parameter that estimates the extent of decrease of J_{ev}(H₂O) for a thick HCl-
364 doped ice film of μm size down to thicknesses of approximately 80 nm.

365 At the start of the evaporation experiment the equilibrium vapor pressure of H₂O,
366 P_{eq}(H₂O), is that of pure ice (Delval et al., 2003; Delval and Rossi, 2004; Pratte et al., 2006)
367 owing to the small values of χ_{HCl}^0 . Raoult’s Law applies to such small values of χ_{HCl} but leads
368 to unmeasurably small deviations from the observed vapor pressure of H₂O which is that of
369 pure ice. In fact, we have never observed an equilibrium vapor pressure that did not
370 correspond to pure ice in the course of the present work that seems to be the consequence of
371 the small average mole fractions of HCl in the H₂O/HCl system. This value of P_{eq}(H₂O) is
372 observed throughout the evaporation up to t_{He} as the film is apparently sufficiently H₂O-rich
373 to support an equilibrium vapour pressure characteristic of pure ice consistent with the
374 published, albeit revised HCl/H₂O phase diagram by Iannarelli and Rossi (2014). In view of
375 the decreasing values of J_{ev}(H₂O) displayed in Figure 2 the equilibrium vapour pressure of
376 pure ice can only be maintained if the condensation rate coefficient k_c for H₂O adsorption
377 decreases to the same extent as J_{ev}(H₂O) in agreement with previous work (Delval et al.,

378 2003; Delval and Rossi, 2004; Pratte et al., 2006) and the concept of microscopic
379 reversibility.

380 Figure 1A displays both the QCMB signal (\square) as well as the corresponding MS signal
381 for evaporating H_2O at $m/e = 18$ (\diamond). Akin to the $\text{HNO}_3/\text{H}_2\text{O}$ system studied previously
382 (Delval and Rossi, 2005) we obtain a perfect match between the two signals for $t < t_D$ whereas
383 for $t > t_D$ there is a significant discrepancy, especially at $t > 300$ s amounting to typically less
384 than a factor of two. Such a disagreement has been noted before for $\text{HNO}_3/\text{H}_2\text{O}$, albeit to a
385 larger extent. The reason for this behaviour of the QCMB signal has not been studied in
386 detail but may well lie in a structural rearrangement of the condensed phase during
387 evaporation that will lead to a change in the calibration factor C_f defined in Table 1 and in
388 Delval and Rossi (2005). In view of the straightforward interpretation of the calibrated MS
389 signal at $m/e = 18$ we have used it for the measurement of $J_{\text{ev}}(\text{H}_2\text{O})$ at $t > t_D$ akin to the
390 previous study on $\text{HNO}_3/\text{H}_2\text{O}$.

391 The accuracy with which both t_{Hb} and t_{He} can be determined depends on the temporal
392 change of the background MS signal for HCl at $m/e = 36$ displayed in Figure 1B following
393 the dosing of the thin ice film under stirred flow conditions. Figure 1B displays the MS signal
394 at $m/e = 36$ as a function of time just before the start of HCl desorption at t_{Hb} that is signalled
395 by an increase in the MS intensity whereas t_{He} corresponds to the return of the HCl signal to
396 the decaying HCl background in comparison to a reference experiment in which the HCl
397 background was monitored as a function of time following the admission of the same HCl
398 dose in the absence of an ice film. We estimate that t_{Hb} is determined to ± 10 s whereas t_{He}
399 may only be estimated to ± 100 s by virtue of the vanishing intensity of the HCl MS signal
400 compared to its slowly decaying background.

401 Previous work has established that the rate of deposition of HCl , R_{HCl} , in the range
402 1×10^{13} to 5×10^{13} molecule s^{-1} for the 0.78 cm^2 surface area of the Si-window leads to the
403 formation of a crystalline HCl hydrate, $\text{HCl} \cdot x\text{H}_2\text{O}$, whereas values outside of this range
404 seemed to favor the formation of an amorphous $\text{HCl}/\text{H}_2\text{O}$ mixture (Delval et al., 2003). The
405 exact nature of this undoubtedly crystalline solid is still unknown. However, IR spectroscopic
406 work on hydroxonium salts of the type $\text{H}_3\text{O}^+\text{X}^-$ suggests that the ν_1 and ν_3 peak positions of
407 the symmetric and antisymmetric O-H stretch vibrations must correspond to a molecular
408 structure in which the distance between the cation and anion is unusually large (Desbat and
409 Huong, 1975; Iannarelli and Rossi, 2016). Recent work has shown that the presence of HCl
410 hexahydrate ($\text{HCl} \cdot 6\text{H}_2\text{O}$) under the present experimental conditions could be safely excluded,
411 however, the FTIR absorption spectrum clearly shows the presence of dissociated HCl within

412 the ice film (Iannarelli and Rossi, 2014). Akin to $\text{HCl}\cdot 6\text{H}_2\text{O}$ that is known to nucleate with
413 difficulty, crystallization of this unknown HCl hydrate seems to occur only under specific
414 conditions of temperature and/or HCl deposition. Owing to the quantitative control of HCl
415 deposition on the ice film in this work we infer the presence of at least two forms of HCl
416 hydrates in the temperature range chosen in analogy to previous work (Delval et al., 2003).

417 We clearly point out that the present work has been performed without simultaneous
418 spectroscopic control of the HCl/ice deposit that would have allowed the identification and/or
419 quantification of the molecular composition of the condensate. Because we lack a
420 spectroscopic probe for the ice film deposited on the QCMB in the present work we are
421 seeking a correlation between the type of HCl/ H_2O deposit, either crystalline or amorphous,
422 and the relevant HCl deposition parameters. Previous work has revealed a distinctly different
423 temporal dependence of $J_{\text{ev}}(\text{H}_2\text{O})$ between the crystalline and amorphous HCl hydrates with
424 the extent of H_2O evaporation from the film, both at low (Delval et al., 2003) and high
425 temporal resolution (Iannarelli and Rossi, 2014).

426 Datasets A and B have been characterized above in terms of a difference in the temporal
427 dependence of $J_{\text{ev}}(\text{H}_2\text{O})$ as a function of increasing χ_{HCl} owing to H_2O evaporation. Taking
428 one example of each set Figure 2 reveals a distinct difference between experiment 7 (set B)
429 and 11 (set A) performed at $T = 195$ and 192 K, respectively, despite comparable HCl
430 deposition parameters (Table 2). At $t > t_{\text{D}}$ $J_{\text{ev}}(\text{H}_2\text{O})$ for experiment 7 decreases at once with
431 χ_{HCl} in contrast to experiment 11 whose $J_{\text{ev}}(\text{H}_2\text{O})$ value gradually starts to decrease at roughly
432 the same value of χ_{HCl} as experiment 7. In addition, in both cases the extent of the decrease of
433 $J_{\text{ev}}(\text{H}_2\text{O})$ is roughly equal between t_{D} and t_{Hb} within less than a factor of two. Set B data are in
434 marked contrast to set A independent of the magnitude of χ_{HCl} which is highlighted by a
435 comparison of experiment 11 (set A) and 4 (set B) at 192 and 190 K, respectively. The abrupt
436 decrease of $J_{\text{ev}}(\text{H}_2\text{O})$ for set B as well as the gradual decline for set A both at t_{D} occur before
437 HCl starts to evaporate from the sample at t_{Hb} and appear therefore to be independent of χ_{HCl}
438 within the range explored in the present work.

439 If we consider the mean value $\langle d_{\text{D}} \rangle$ for data set B (Figures 4 and A2) we find $549.0 \pm$
440 120.0 nm compared to the $1'500$ nm or so original ice thickness which corresponds to
441 approximately 8.5×10^{17} molecules of H_2O spread out over 0.50 cm^2 . These H_2O molecules
442 are impacted by the presence of HCl to some extent because $J_{\text{ev}}(\text{H}_2\text{O})$ is slowed down
443 significantly compared to pure ice. Previous results (Delval et al., 2003) on the deposition of
444 HCl on ice under conditions where the presence of an as yet unidentified crystalline hydrate
445 $\text{HCl}\cdot x\text{H}_2\text{O}$ was confirmed by FTIR absorption led to the conclusion that on average the

446 amount of “trapped” H₂O within d_D corresponded to 1.2x10¹⁸ molecules starting with an
447 original 1 μm thick ice film that was subsequently doped with HCl. This quantity of H₂O,
448 when scaled from the 0.78 cm² area of the Si-window used for FTIR absorption to the area of
449 0.5 cm² of the QCM leads to 7.7x10¹⁷ H₂O that is in satisfactory agreement with the present
450 measurement of d_D or 8.5x10¹⁷ H₂O in the present work. We may add that the previous value
451 of 1.2x10¹⁸ H₂O from the work of Delval et al. (2003) corresponding to d_D obtained in that
452 work has been derived using HeNe interferometry which is a crude method for measuring the
453 film thickness.

454 Specifically, considering the low value of d_D of experiments 1 and 10 (Table 2, Figure
455 4) we may define the behaviour of these condensates as “ice-like” because roughly 80% of the
456 ice sample of roughly 1.5 μm thickness has evaporated at J_{ev}(H₂O) of pure H₂O ice before it
457 slows down. This decrease of J_{ev}(H₂O) is a kinetic effect and acts on both the rate of
458 evaporation as well as on the mass accommodation coefficient, the ratio of which remains
459 constant because the characteristic vapor pressure of pure ice is maintained until t = t_{He} when
460 the sample runs out of H₂O and HCl. For sample 1 this conclusion is not too surprising owing
461 to its extremely low HCl dose of 0.8 formal HCl monolayers. Sample 10 in comparison with
462 the other members of data set A allows us to conclude that d_D is proportional to T_{ice} for data
463 set A. Low temperatures prevent rapid diffusion of HCl into the bulk of the ice film which
464 leaves the majority of the total mass of the thin film deposited void of any HCl. Therefore, a
465 large fraction of the total mass of the thin film deposit evaporates at values of J_{ev}(H₂O)
466 characteristic of pure ice before it decreases to lower values when the presence of HCl slows
467 down J_{ev}(H₂O). Although our experiment does not reveal the location of the thin layer of HCl-
468 contaminated ice, plausibility suggests that it is located on top of the ice film at the gas-
469 condensed interface. The corollary of this is that it is impossible to “cap” a pure ice sample
470 with a thin layer of an atmospheric condensable gas of lower vapor pressure in the hope to
471 lower the vapour pressure of the condensate or slow down H₂O evaporation. This capping has
472 been attempted many times, and examples abound. However, all attempts to lower the ice
473 vapor pressure of the condensate using low amounts of polar contaminants of ice, such as
474 HNO₃, HCl or HBr have proven futile to date (Biermann et al., 1998).

475 The other members of data set A are examples (experiments 2, 9, 11) with high values
476 of d_D at higher temperatures and higher HCl doses (Table 2). Because of higher presumed
477 interfacial HCl concentrations these samples experience a decrease in J_{ev}(H₂O) owing to rapid
478 diffusion of HCl into ice that affects the kinetics of evaporation to some depths of the ice film

479 corresponding to higher values of d_D . Both high HCl doses and high temperatures favor HCl
480 contamination of deeper layers of the HCl film, hence high values of d_D .

481 Tentatively, we assign a crystalline, yet unknown molecular structure and stoichiometry
482 to samples A in contrast to samples of dataset B that we identify with an amorphous structure
483 in terms of a liquid HCl/H₂O mixture of variable composition. The main argument in favour
484 of this assignment comes from recent kinetic work performed by Iannarelli and Rossi (2016a)
485 who show that both $J_{ev}(H_2O)$ as well as the corresponding mass accommodation coefficient or
486 the adsorption rate coefficient for H₂O adsorption is highly scattered for crystalline HCl
487 hexahydrate whereas the amorphous mixture shows a significantly smaller scatter of the
488 experimental and thermodynamic values (Iannarelli and Rossi, 2014). Figure A3 and A4 in
489 the Appendix show this substantial difference in experimental scatter for the amorphous
490 HCl/H₂O mixture (Figure A3) compared to crystalline HCl hexahydrate (Figure A4).

491 Figure 3 displays the range parameter $r^{b/e}$ as a function of R_{HCl} for all data displayed in
492 Table 2. It is noteworthy that $r^{b/e}$ is in the range 20 to 27 for set B experiments 3, 4, 7 and 8
493 compared to set A data that seem to be scattered throughout the range. Members of data set B
494 show a common average range for both d_D and $r^{b/e}$ which is the reason we tentatively assign
495 these structures to amorphous liquid mixtures of high viscosity at the prevailing temperatures.

496 In conclusion, we take the simultaneous occurrence of the restricted range of the
497 measured remaining thickness of ice $d_D = 549.0 \pm 120.0$ nm together with a similarly
498 restricted range of $r^{b/e}$ between 20 and 27 as well as the substantial overlap in R_{HCl} between
499 the present and previous work (Delval et al., 2003) as an indication that set B evaporation
500 experiments imply the presence of an amorphous HCl/H₂O mixture. In contrast, the scatter of
501 the set A data across the range of $r^{b/e}$ and d_D values suggests the presence of an as yet
502 unidentified crystalline HCl hydrate. If, and only if the HCl deposition conditions rapidly
503 establish thermodynamic equilibrium, experiment 2 (low HCl flow rate) lies in the “ice”
504 region in the temperature interval 192-210 K whereas experiment 11 (high HCl flow rate)
505 should access crystalline HCl hexahydrate at 192 K but not at 210 K according to the revised
506 HCl/H₂O phase diagram of Iannarelli and Rossi (2014). It remains to be seen whether or not
507 the published FTIR absorption spectrum in Delval et al (2003) turns out to be identical to the
508 expected crystalline HCl hexahydrate invoked as condensate in set A molecules, similar HCl
509 deposition parameters notwithstanding. This proposal awaits further confirmation from FTIR
510 spectroscopic work that will be combined in the future with the QCM measurement. At this
511 point we reiterate our earlier statement that T_{ice} , R_{HCl} , N^{dep}_{HCl} do apparently not control
512 $J_{ev}(H_2O)$ of both datasets..

513 5. ATMOSPHERIC IMPLICATIONS

514

515 The evaporation range parameter $r^{b/e}$ may be used to quantitatively evaluate the upper
516 limit of the evaporative lifetime extension of thin ice films under conditions of H₂O vapor
517 subsaturation. In the interest of applying the data of the present work to atmospheric
518 conditions we make the assumption that typical atmospheric Cirrus cloud particles of several
519 μm diameter may be approximated by macroscopic thin films used to obtain the present data.
520 The time t_{ev} in seconds to complete evaporation of an ice particle of radius r at a given relative
521 humidity (rh) is given in equation 1 (Chiesa and Rossi, 2013; Iannarelli and Rossi, 2016a):

$$522 \quad t_{ev} = \frac{\left(\frac{\rho N_L}{M}\right)^{2/3} \left(\frac{r}{a}\right)}{J_{ev} (1 - rh)} \quad (1)$$

523 where ρ is the density of ice (0.916 and 0.925 g cm⁻³ at 273 and 173 K, respectively),
524 $M = 18 \text{ g mol}^{-1}$ for H₂O, r and a are the ice particle radius and the distance between two
525 molecular layers in H₂O(ice), respectively (Iannarelli and Rossi, 2016a). Equation 1 is based
526 on a simple layer-by-layer evaporation model of H₂O(ice) from a spherical ice particle
527 following a zero-order rate law for J_{ev} or a first order rate law for its inverse, namely H₂O
528 adsorption or condensation. For a 10 μm diameter ice particle approximated by thin film
529 experiment 1 (Table 2) at $rh = 80\%$, $T = 192 \text{ K}$, $J_{ev} = 3 \times 10^{16} \text{ molecule s}^{-1} \text{ cm}^{-2}$ (Petrenko and
530 Whitworth, 1999) and $a = 4 \times 10^{-8} \text{ cm}$ we obtain $t_{ev} = 2050 \text{ s}$ or 34 min. This is the value for a
531 pure ice particle as $J_{ev}(\text{H}_2\text{O})$ for pure ice has been used at the outset of the evaporation
532 experiment and is a lower limit to the true evaporation time owing to the competition of mass
533 transfer and heterogeneous chemistry (Seinfeld and Pandis, 1998). Using $r^{b/e} = 43$ for
534 experiment 1 t_{ev} is calculated to be 15 minutes and 24 hours for a 100 nm and 10 μm diameter
535 particle, respectively, whereas the evaporative lifetime of an analogous pure ice particle
536 would be only 21 s for the 100 nm diameter pure ice particle. Cirrus ice particles are
537 frequently in the lower tens of μm size range resulting in a longer evaporation time
538 considering that the simple evaporation model scales linearly with the radius of the ice
539 particle. In conclusion we may state that, owing to the lifetime extension of ice particles
540 contaminated by HCl, HNO₃ or other volatile atmospheric trace gases such as HOCl, HOBr or
541 HONO, small particles may have a chance to survive subsaturated regions of the atmosphere
542 so as to function as cloud condensation or ice nuclei for the following cloud cycle (Delval and
543 Rossi, 2004; 2005; Pratte et al., 2006).

544 We would like to stress, that the variable $r^{b/e}$ factor displayed in Table 2 leads to a
545 significant increase in the evaporative lifetime of a contaminated ice particle and amounts to a
546 **kinetic effect** that does not affect the equilibrium vapor pressure of the ice particle in
547 question: it is that of pure ice from the start of the evaporation experiment to $t = t_{He}$ and
548 therefore affects both the rate of evaporation and accommodation equally. However, in cases
549 the sample has lost most of its mass the vapor pressure decreases and becomes somewhat
550 uncertain. In the present case the above statement is correct for $t = t_{Hb}$, that is before halogen
551 evaporation. Of note is the fact, that the accommodation coefficient α is frequently less than
552 unity, in contrast to what is often assumed, which will lower the rate of evaporation for pure
553 ice, hence increase the evaporative lifetime of pure ice particles for $T \geq 180$ K as proposed in
554 previous work (Delval and Rossi, 2004; 2005; Pratte et al., 2006).

555 As a token example of potential atmospheric importance of the measured evaporative
556 lifetimes of ice particles laced with condensable atmospheric trace gases we may take the
557 formation, persistence and evaporation of contrails and Cirrus clouds in the UT/LS. These are
558 ice clouds forming on non-volatile ice nuclei at the corresponding temperature and relative
559 humidity conditions and also frequently serve as reaction sites for heterogeneous atmospheric
560 reactions in connection with ozone depletion and chlorine activation chemistry in the LS.
561 Under certain conditions, Schumann and coworkers used the concept of the increase of the
562 evaporative lifetimes of contaminated ice particles in aviation contrails occurring mostly in
563 the UT, but sometimes also in the LS, in order to explain the persistence of ice clouds below
564 ice saturation conditions up to a certain time duration. Ice clouds have a significant radiative
565 forcing effect that is of interest in evaluating the climate forcing of high-flying aircraft in
566 future aviation scenarios (Lewellen, 2014; Schumann et al, 2017; 2017a). However, the
567 results of the present work show that the rate of evaporation of ice films doped with small
568 amounts of acidic trace gases significantly slows down in a complex manner over the
569 evaporation history of the film or particle, and that the application of equation (1) to
570 atmospheric situations should be carried out with caution.

571

572 CONCLUSIONS

573

574 Despite the scatter of the values of $r^{b/e}$ and d_D in dataset A displayed in Figures 3 and 4
575 and the apparent lack of influence of the deposition parameters (T_{ice} , R_{HCl} , N^{dep}_{HCl}) on
576 $J_{ev}(H_2O)$ we may state several key points from the present work:

- 577 (a) We observe two types of behaviour, both complex, as far as the temporal change of
578 $J_{ev}(H_2O)$ with on-going evaporation of H_2O from a HCl/H_2O condensate is
579 concerned. We have named it sets A and B that represent limiting behaviour as not
580 all performed experiments fit into this scheme.
- 581 (b) At low temperature or low dose of deposited HCl (N_{HCl}^{dep}) set A samples, especially
582 samples 1 and 10, reveal an “ice-like” behaviour that corresponds to a low value of
583 d_D . This means that the HCl/H_2O condensate evaporates a large fraction of the
584 sample thickness at a value of $J_{ev}(H_2O)$ characteristic of pure ice before slowing
585 down at increasing mole fraction of HCl upon H_2O evaporation. This corresponds to
586 a two-phase system consisting of a major ice-like and a minor HCl/H_2O phase
587 having both significantly different values of $J_{ev}(H_2O)$.
- 588 (c) High values of d_D are observed at high T_{ice} or N_{HCl}^{dep} values for set A samples. This
589 means that the sample evaporates H_2O at $J_{ev}(H_2O)$ characteristic of pure ice for a
590 relatively short time of its evaporation history because the quantity of HCl is
591 sufficient to decrease $J_{ev}(H_2O)$ already at high values of d_D by rapidly diffusing to
592 deeper layers of the ice film. An equivalent way of expressing the point would be to
593 state that d_D which is an indicator of the total mass of the ice film, is proportional to
594 T_{ice} for Set A.
- 595 (d) Set A samples generally show scattered values of both d_D and $r^{b/e}$ values that we
596 attribute to the existence of a two-phase binary system, namely a pure ice and a
597 crystalline HCl hydrate phase of as yet unknown stoichiometry $HCl \cdot xH_2O$, but
598 probably HCl Hexahydrate. At first the pure ice phase starts to evaporate as a whole
599 for a fairly long time at characteristic values of $J_{ev}(H_2O)$ until the pure ice phase has
600 disappeared, followed by the crystalline HCl/H_2O phase at a lower rate of $J_{ev}(H_2O)$
601 to attain the characteristic value for the evaporation of the crystalline $HCl \cdot xH_2O$
602 phase.
- 603 (e) Set B samples are tentatively identified as single phase binary amorphous mixtures
604 of HCl/H_2O whose kinetic properties are uniform, thus fairly independent of the HCl
605 concentration at the gas-condensed phase interface. The observation of a medium
606 size average value for both $r^{b/e}$ and d_D is consistent with these observations and
607 manifests itself as a continuous, yet gradual decrease of $J_{ev}(H_2O)$ with increasing
608 χ_{HCl} . It is in distinct contrast to Set A where $J_{ev}(H_2O)$ values are those of pure ice
609 until the ice phase has completely evaporated followed by a gradual decline of
610 $J_{ev}(H_2O)$ when the crystalline HCl hydrate starts to decompose.

611 (f) It must be recalled that the vapour pressure of H₂O remained that of pure ice during
612 most of the thickness of the H₂O/HCl condensate down to approximately 80 nm at
613 which point we halted the evaporation experiment. This result is expected based on
614 Raoult's law owing to the small average HCl mole fractions in doped ice used in the
615 present work: It would make the decrease of the H₂O saturation vapour pressure
616 unmeasurably small. The present results therefore primarily address the **kinetics** of
617 H₂O evaporation which changes with the total mass of the thin film condensate and
618 the concomitant increase in HCl concentration and/or mole fraction.

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620

621 **References**

622

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802 The authors declare no competing interests regarding the present work.

803

804 **Acknowledgement**

805 We sincerely thank Dr. R. Iannarelli for Figures A3 and A4 displayed in the Appendix. We
806 also would like to thank the Swiss National Science Foundation (SNSF) for unfailing support
807 over the years. This work has been performed under SNSF grants no. 20-65299.01 and
808 200020-105471.

809

Table 1: Hardware parameters of both cryogenic sample supports

	Si Optical Window	QCM	
Reactor temperature T_r [K]	320		
Reactor volume V_r [cm ³]	2350		
Conversion factor (1/RT) Conv [molec cm ⁻³ Torr ⁻¹] with R=62398 [Torr cm ³ mol ⁻¹ K ⁻¹]	$3.0 \cdot 10^{16}$ ⁽¹⁾		
Sample surface area [cm ²]	0.78	0.50	
H ₂ O collision frequency with ice sample ω_{H_2O} [s ⁻¹]	5.08	3.26	
H ₂ O effusion rate constant of calibrated leak $k_{esc}(H_2O)$ [s ⁻¹]	0.064		
MS calibration factor for H ₂ O (m/z=18, Stirred Flow) C_{18}^{S-Flow} [molec s ⁻¹ A ⁻¹]	$2.4 \cdot 10^{24}$		
MS calibration factor for H ₂ O (m/z=18, Dynamic) C_{18}^{dyn} [molec s ⁻¹ A ⁻¹]	$1.7 \cdot 10^{25}$		
HCl collision frequency with ice sample ω_{HCl} [s ⁻¹]	3.59	2.31	
HCl effusion rate constant of calibrated leak $k_{esc}(HCl)$ [s ⁻¹]	0.047		
MS calibration factor for HCl (m/z=36, Stirred Flow) C_{36}^{S-Flow} [molec s ⁻¹ A ⁻¹]	$3.9 \cdot 10^{24}$		
MS calibration factor for HCl (m/z=36, Dynamic) C_{36}^{dyn} [molec s ⁻¹ A ⁻¹]	$6.3 \cdot 10^{24}$		
Calculated escape orifice area A_{esc} [mm ²]	1.0		
	$d = 10^4 \text{ \AA}$ or 1.0 \mu m for O.D.= 1.08 [A] ⁽²⁾ at 3260 cm^{-1}		Calibration Factor
		Temperature [K]	ratio ⁽³⁾
		170	9.0
		180	8.0
		190	7.8
		193	6.0
		205	2.0
		208	1.9

¹ Wall temperature of the reactor at T = 320 K

² See (Delval et al., 2003)

³ Corresponds to the ratio between the true number of molecules present on the QCM support and the number of molecules displayed by the IC5 controller (Delval et al., 2004)

Table 2 : Representative experimental results for the kinetics of H₂O evaporation in the presence of HCl for increasing HCl deposition temperatures at given rates of deposition R_{HCl} and doses of HCl N_{HCl}^{dep}. In the first column the number refers to the corresponding experiment and identifies the data displayed in Figure 2

Experiment number	T _{ice} [K]	d ₀ [Å]	N _{H₂O} ⁰ [molec]	R _{HCl} [molec s ⁻¹]	t _{dep} [s]	N _{HCl} ^{dep} [molec]	HCl ML	N _{HCl} ^{evap} [molec]	χ _{HCl} ⁰	d _D [Å]	J _{ev} ^b [molec cm ⁻² s ⁻¹]	J _{ev} ^e [molec cm ⁻² s ⁻¹]	r ^{b/e}
10	174	15230	2.4 · 10 ¹⁸	6.4 · 10 ¹²	94	6.0 · 10 ¹⁴	4.8	4.7 · 10 ¹⁴	2.5 · 10 ⁻⁴	2733	1.9 · 10 ¹⁵	4.4 · 10 ¹⁴	4.3
5	188	13318	2.0 · 10 ¹⁸	1.3 · 10 ¹³	66	8.7 · 10 ¹⁴	7.0	8.9 · 10 ¹⁴	4.4 · 10 ⁻⁴	4540	1.2 · 10 ¹⁶	3.9 · 10 ¹⁵	3.1
4	190	14016	2.1 · 10 ¹⁸	4.2 · 10 ¹³	126	5.4 · 10 ¹⁵	43.2	3.6 · 10 ¹⁵	2.6 · 10 ⁻³	6360	2.9 · 10 ¹⁶	1.4 · 10 ¹⁵	20.7
6	190	13886	2.1 · 10 ¹⁸	3.9 · 10 ¹³	56	2.2 · 10 ¹⁵	17.6	1.8 · 10 ¹⁵	1.0 · 10 ⁻³	12861	3.4 · 10 ¹⁶	1.7 · 10 ¹⁶	2.0
1	192	14926	2.3 · 10 ¹⁸	3.1 · 10 ¹²	36	1.0 · 10 ¹⁴	0.8	1.8 · 10 ¹⁴	4.3 · 10 ⁻⁵	2823	2.9 · 10 ¹⁶	7.1 · 10 ¹⁴	40.8
2	192	14682	2.3 · 10 ¹⁸	8.0 · 10 ¹¹	356	2.6 · 10 ¹⁴	2.1	1.6 · 10 ¹⁴	1.1 · 10 ⁻⁴	6817	3.2 · 10 ¹⁶	6.5 · 10 ¹⁴	49.2
11	192	14420	2.2 · 10 ¹⁸	5.4 · 10 ¹²	108	5.4 · 10 ¹⁴	4.3	6.8 · 10 ¹⁴	2.4 · 10 ⁻⁴	7717	4.0 · 10 ¹⁶	7.9 · 10 ¹⁴	50.6
3	193	14423	2.2 · 10 ¹⁸	3.5 · 10 ¹²	220	7.0 · 10 ¹⁴	5.6	8.1 · 10 ¹⁴	3.2 · 10 ⁻⁴	5659	4.9 · 10 ¹⁶	1.8 · 10 ¹⁵	27.2
7	195	12614	1.9 · 10 ¹⁸	4.3 · 10 ¹²	45	1.9 · 10 ¹⁴	1.5	1.8 · 10 ¹⁴	1.0 · 10 ⁻⁴	5325	4.6 · 10 ¹⁶	2.0 · 10 ¹⁵	23.0
8	205	13505	2.1 · 10 ¹⁸	1.6 · 10 ¹³	36	5.9 · 10 ¹⁴	4.7	3.0 · 10 ¹⁴	2.8 · 10 ⁻⁴	4607	2.0 · 10 ¹⁷	1.0 · 10 ¹⁶	20.0
9	210	13134	2.0 · 10 ¹⁸	3.5 · 10 ¹²	84	3.0 · 10 ¹⁴	2.4	1.9 · 10 ¹⁴	1.5 · 10 ⁻⁴	12136	3.0 · 10 ¹⁷	1.8 · 10 ¹⁶	16.7

Figure Captions

Figure 1: Typical experimental protocol of the evaporation at 192 K of an approximately 1.2 μm thick ice film doped with $5.4 \cdot 10^{14}$ molecules of HCl. This illustration corresponds to experiment 11 of Table 2. (\circ): ice thickness monitored by QCM (\AA), (\square): "apparent" H_2O evaporative flux, $J^{\text{QCM}}_{\text{ev}}$, monitored by QCM ($\text{molecule cm}^{-2} \text{s}^{-1}$), (+): I^{18} MS signal for H_2O , (\times): I^{36} MS signal for HCl (A), (\diamond): J^{18}_{ev} evaporative flux calculated from I^{18} ($\text{molecule cm}^{-2} \text{s}^{-1}$), (Δ): $\text{Int}(J^{18}_{\text{ev}})$ time integral of J^{18}_{ev} (molecule cm^{-2}).

Figure 2: Change of the evaporative flux $J_{\text{ev}}(\text{H}_2\text{O})$ as a function of the HCl mole fraction (χ_{HCl}) for the cases presented in Table 2 color-coded according to the corresponding experiment number in Table 2. The colored and circled numbers on axis "b" (left) correspond to $J_{\text{ev}}(\text{H}_2\text{O})$ of pure ice before HCl deposition, the ones on axis "e" (right) are $J_{\text{ev}}(\text{H}_2\text{O})$ at $t = t_{\text{He}}$ at the end of HCl evaporation. The colored circles in the data field mark the value of $J_{\text{ev}}(\text{H}_2\text{O})$ after HCl deposition at $t = t_0$ and are equal to $J_{\text{ev}}(\text{H}_2\text{O})$ of pure ice. The start of any particular $J_{\text{ev}}(\text{H}_2\text{O})$ curve as a continuous solid (bold) line occurs at $t = t_{\text{D}}$ at 85% of $J_{\text{ev}}(\text{H}_2\text{O})$ at $t = 0$ (pure ice value, colored circle or circled number on axis "b" to the left) and ends at t_{Hb} , the beginning of HCl evaporation as displayed in Figure 1B.

Figure 3: Synopsis of the dependence of the evaporation range parameter $r^{\text{b/e}}$ on the rate of deposition R_{HCl} of HCl for temperatures between 188 and 210 K. Each point is marked with the total number of HCl molecules (N_{HCl}) deposited on the ice film, the temperature of the ice film at HCl deposition and the experiment number (bold) referring to Table 2. The hashed area encompasses $r^{\text{b/e}}$ values for dataset B (experiments 3,4,7,8). The color code goes from low (blue) over medium (green) to high (red) temperatures.

Figure 4: Synopsis of the dependence of d_{D} on the rate of deposition R_{HCl} of HCl for temperatures between 188 and 210 K. Each point is marked with the total number of HCl molecules (N_{HCl}) deposited on the ice film, the temperature of the ice film at HCl deposition and the experiment number (bold font) referring to Table 2. The hashed area encompasses d_{D} values for dataset B (experiments 3,4,7,8). The color code goes from low (blue) over medium (green) to high (red) temperatures.

Figure 1

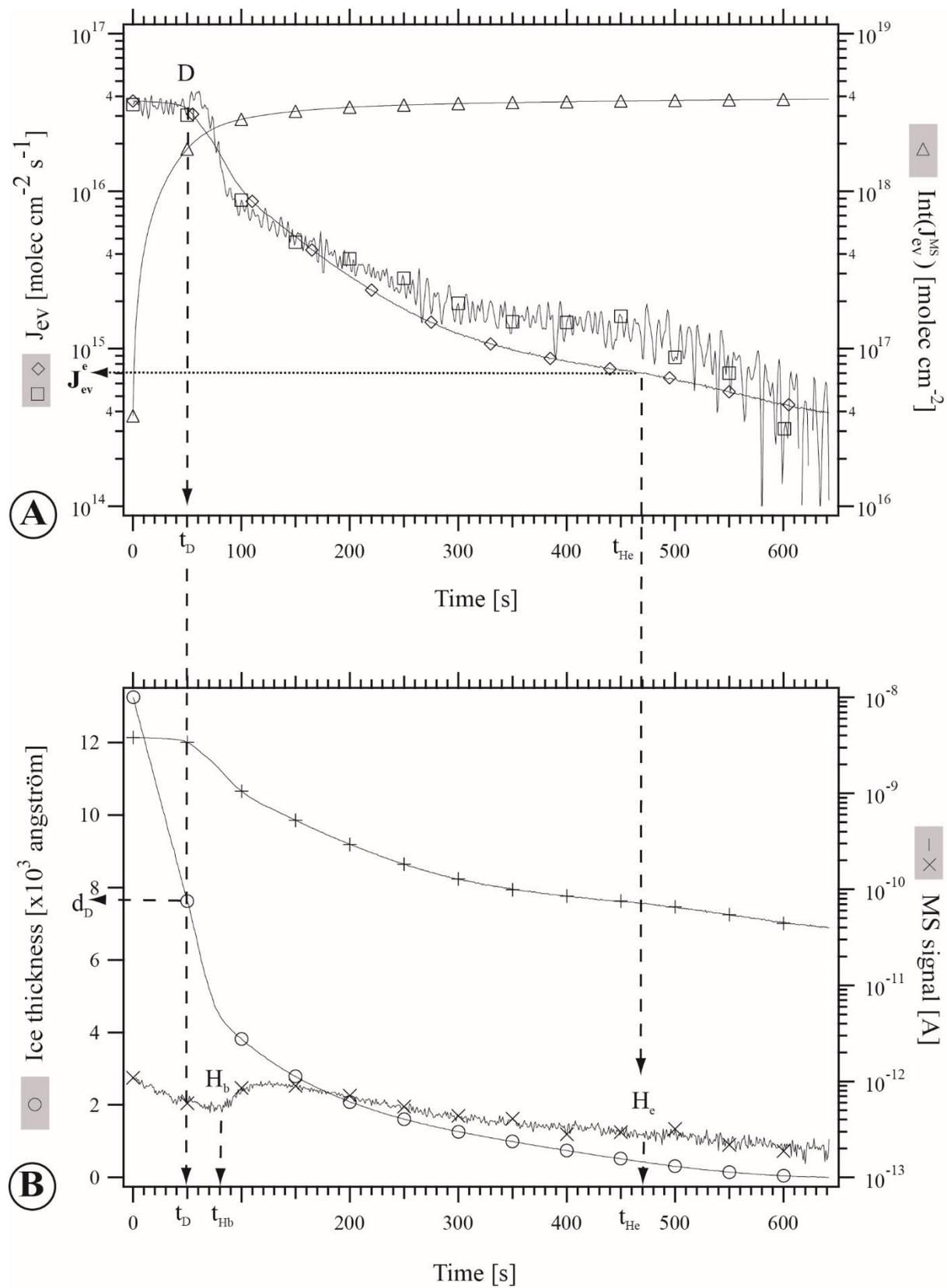


Figure 2

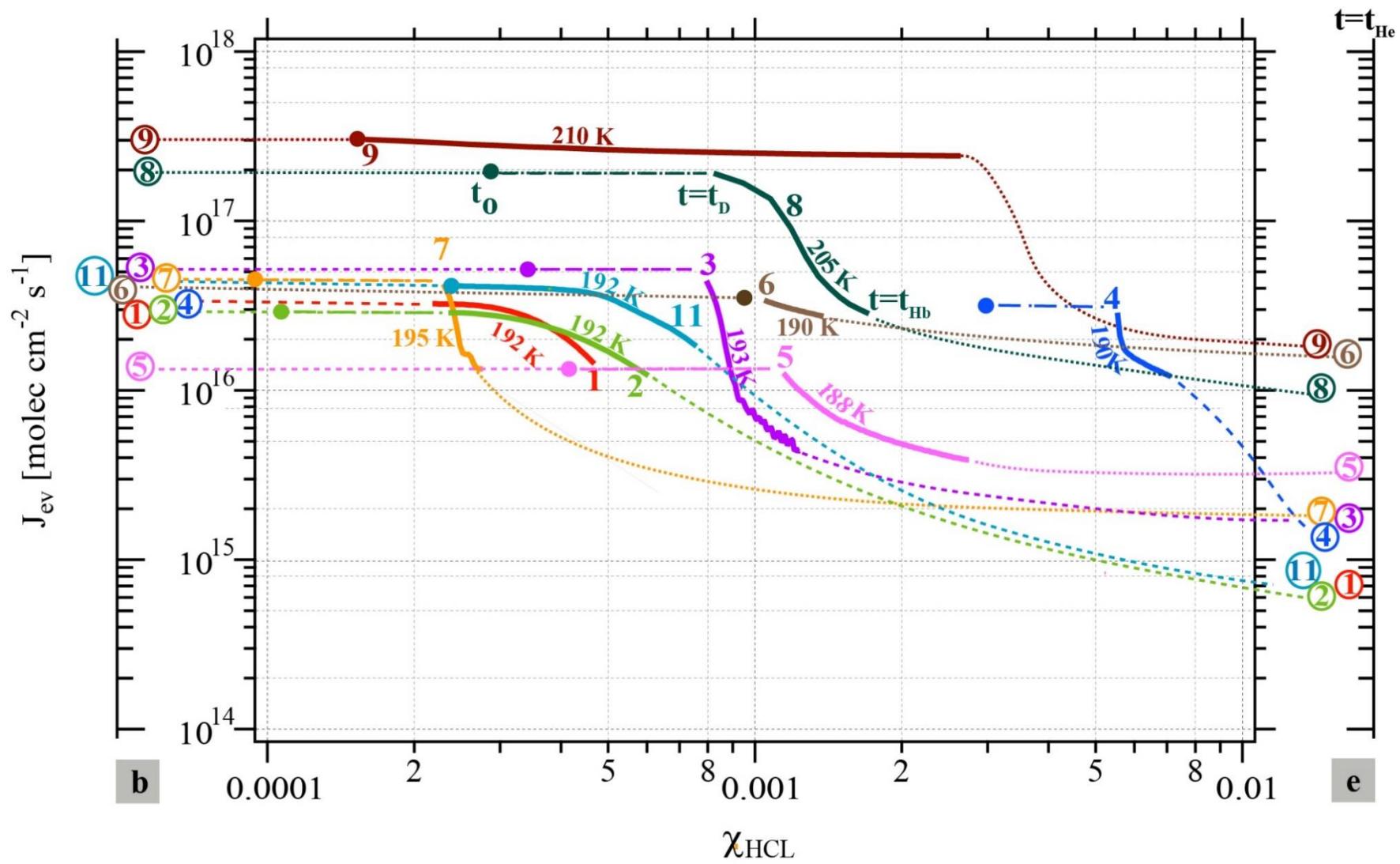


Figure 3

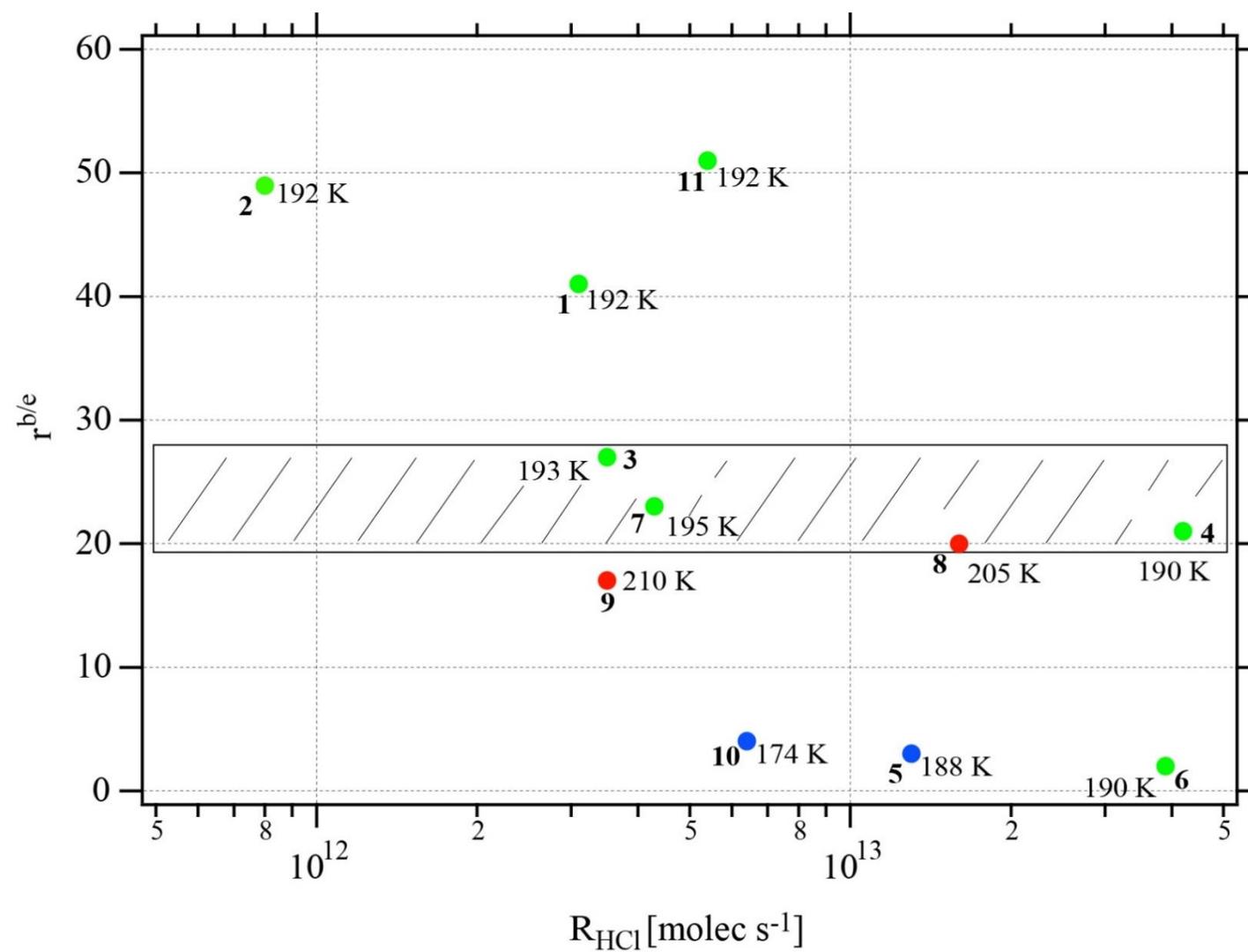
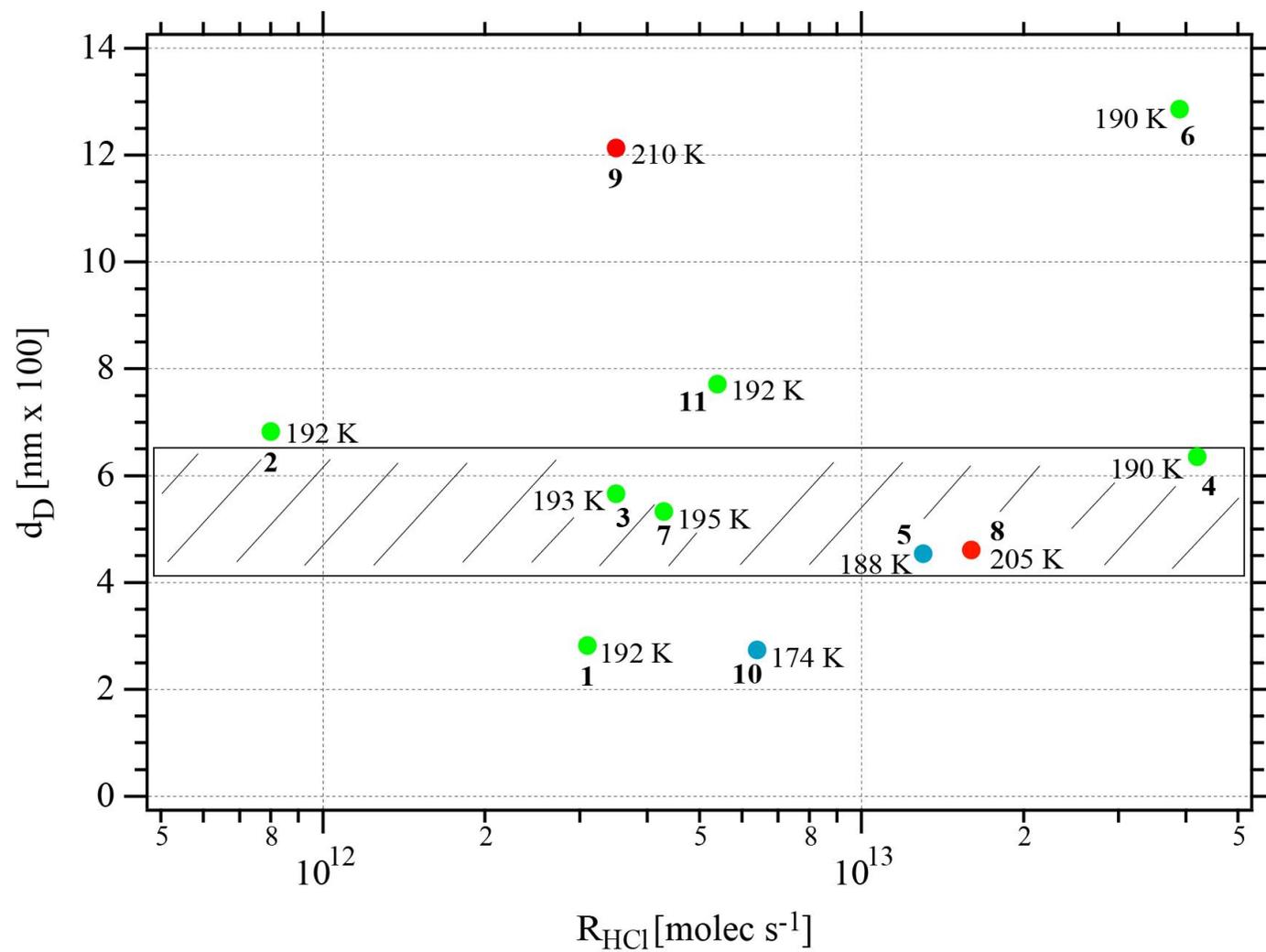


Figure 4



APPENDIX A: Figures A1, A2, A3 and A4, Table A1

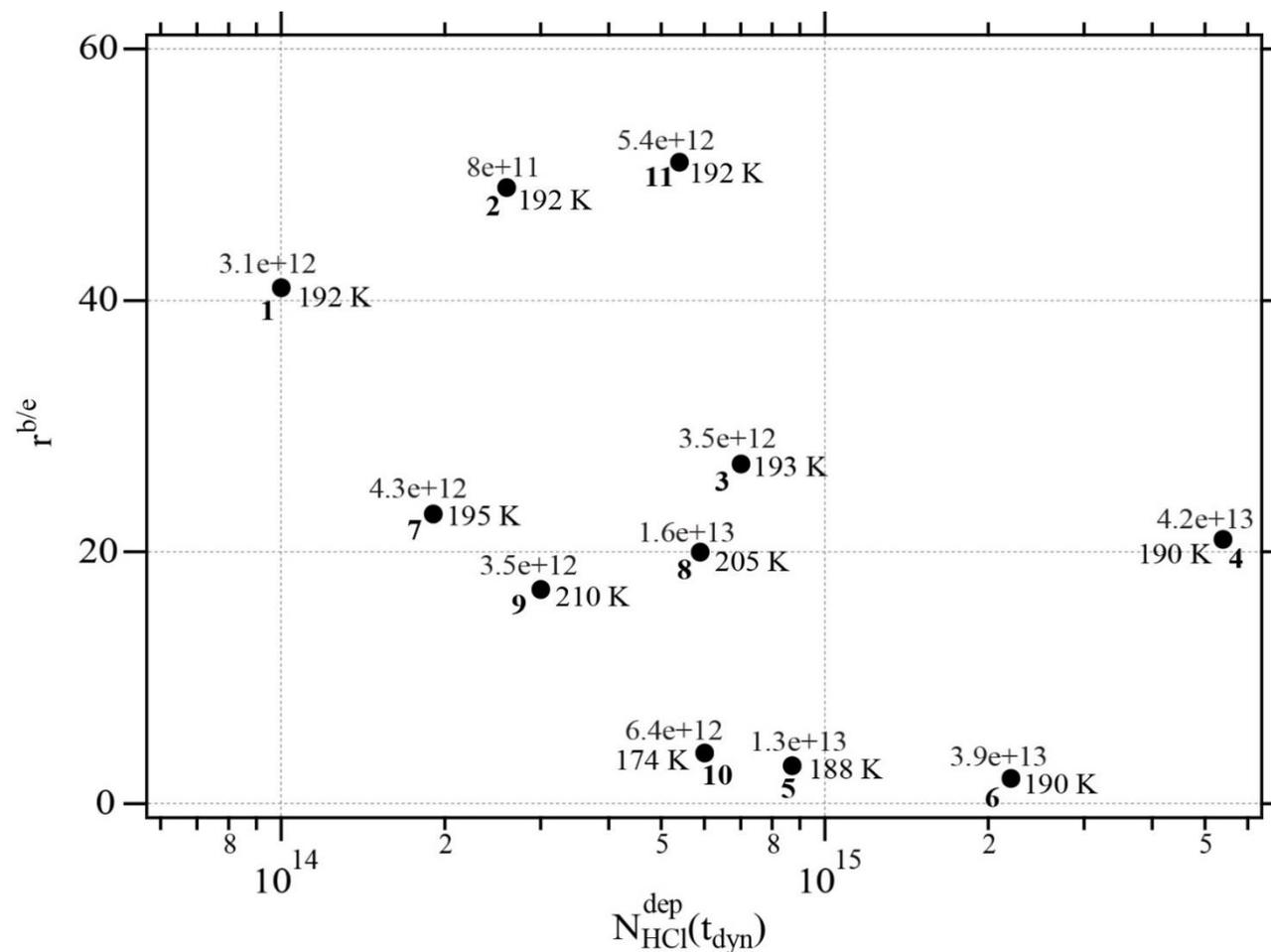


Figure A1: Synopsis of the dependence of the evaporation range parameter, $r^{b/e}$, on the number of adsorbed HCl, N_{HCl}^{dep} , adsorbed on ice for temperatures between 188 and 210 K. each point is marked with the deposition rate of HCl molecules in molec s^{-1} on the ice film, the temperature of the ice film and the experiment running number (bold) referring to Table 2.

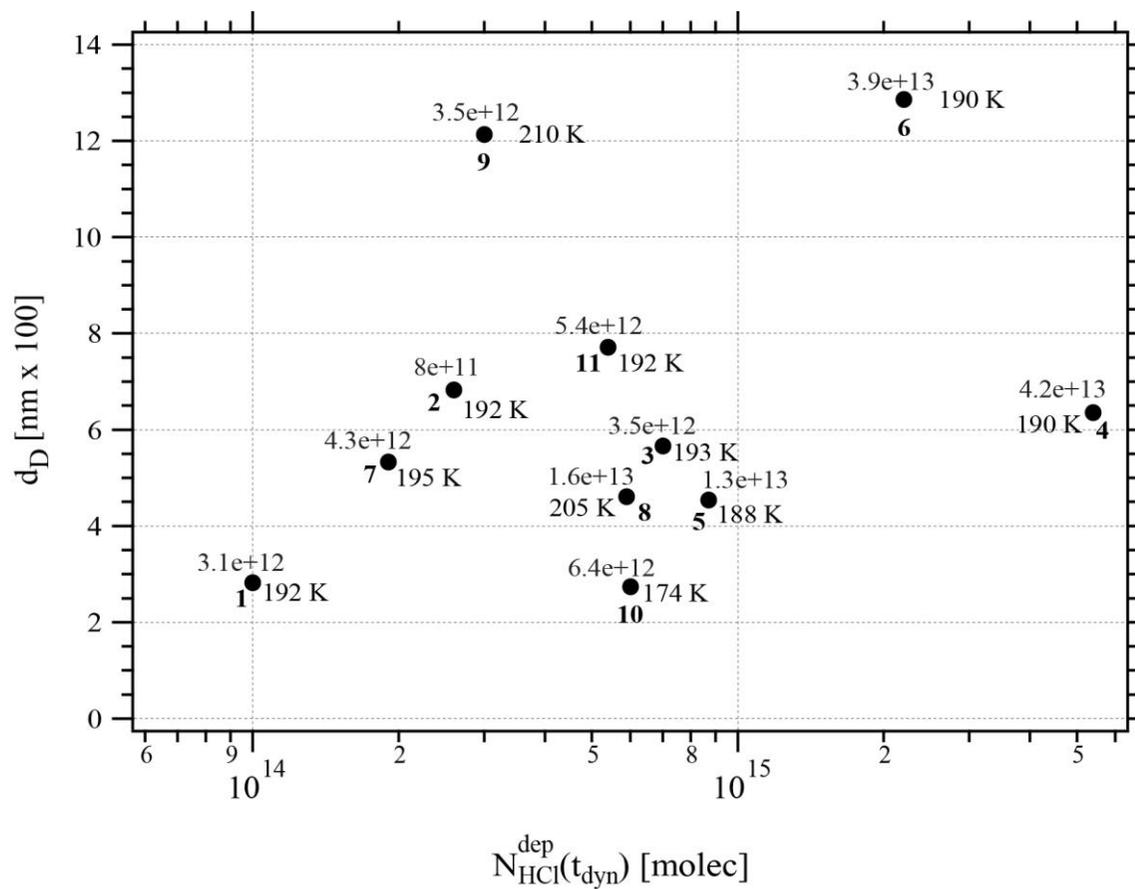


Figure A2: Synopsis of the dependence of the remaining thickness d_D on the number of adsorbed HCl, $N_{\text{HCl}}^{\text{dep}}$, dispensed on ice for temperatures between 188 and 210 K. Each point is marked with the deposition rate of HCl in molec s^{-1} on the ice film, the temperature of the ice film and the experiment running number (bold) referring to Table 2.

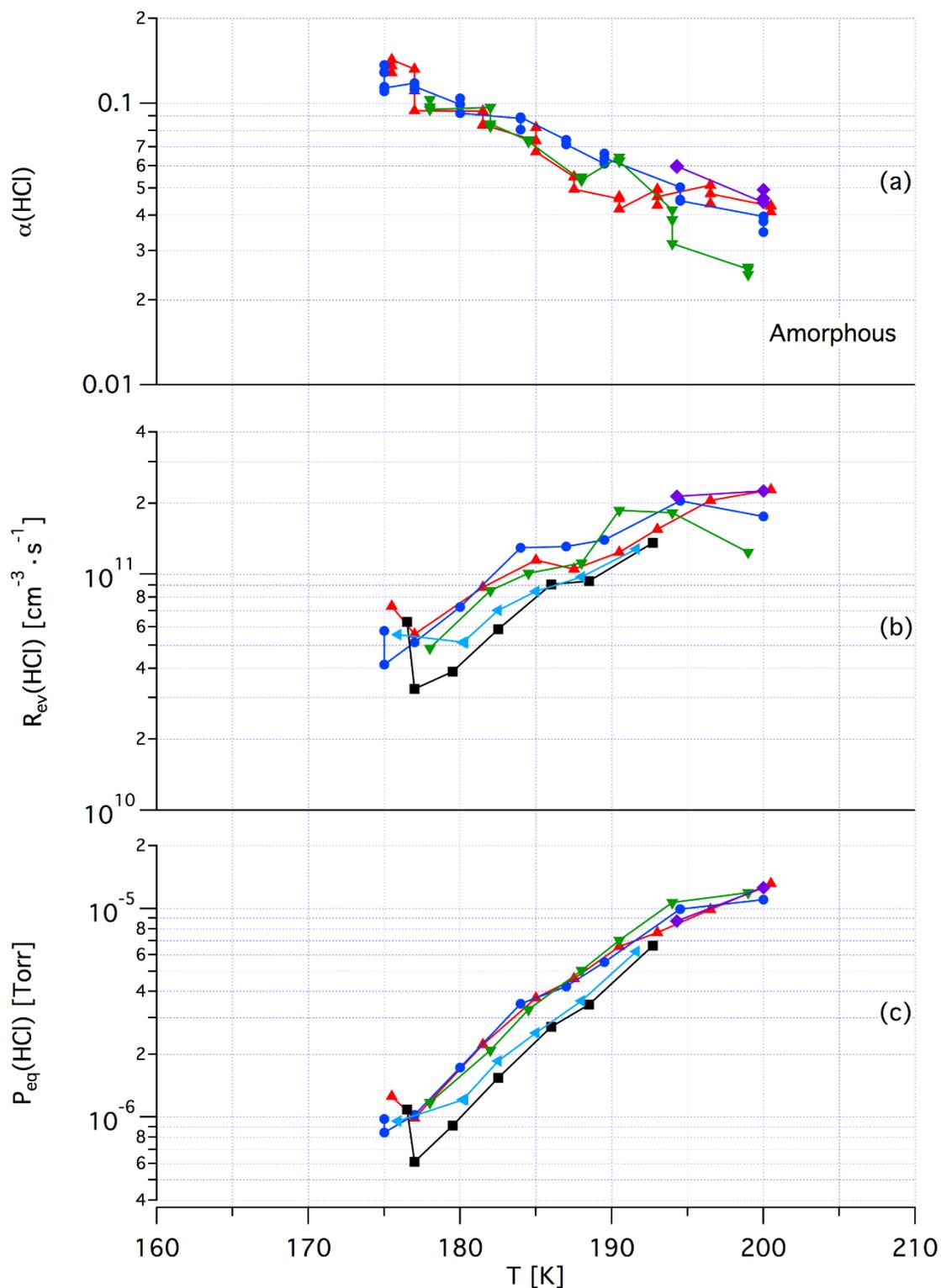


Figure A3: Synopsis of kinetic and thermodynamic results for an amorphous $\text{H}_2\text{O}/\text{HCl}$ mixture using HCl as a probe gas. The symbols/colors used correspond to different experimental runs and the graphs show the scatter of the individual measurements within a series. Original data are published in Iannarelli, R. and Rossi, M.J. *Atmos. Chem. Phys.* 14, 5183–5204, 2014.

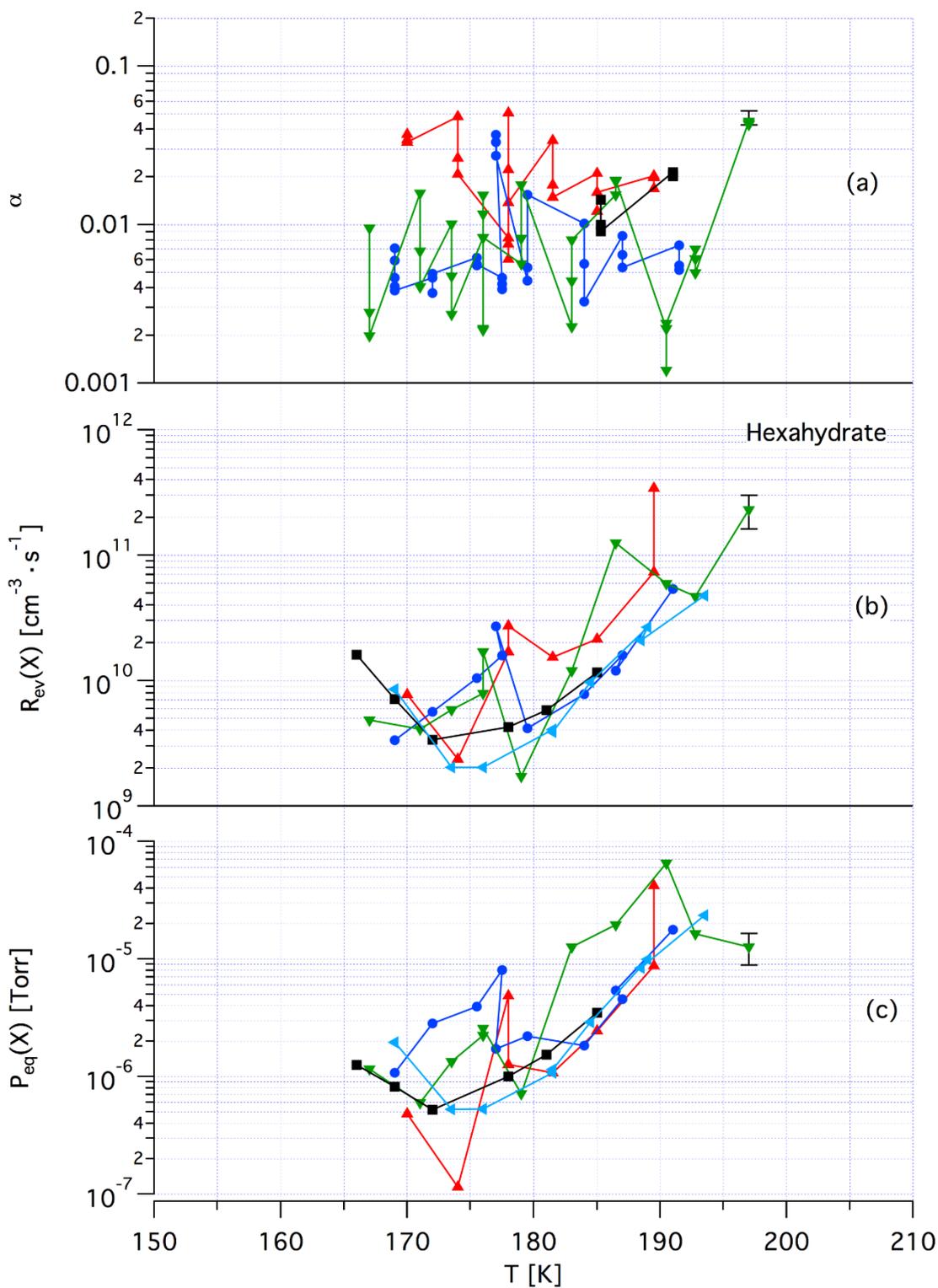


Figure A4: Synopsis of kinetic and thermodynamic results for crystalline HCl hexahydrate (HH) using $X = \text{HCl}$ as a probe gas. The symbols/colors used correspond to different experimental runs and the graphs show the scatter of the individual measurements within a series. Original data are published in Iannarelli, R. and Rossi, M.J. *Atmos. Chem. Phys.* 14, 5183–5204, 2014.

Table A1: Brief Summary of the Amount of a Molecular Monolayer (Coverage) of HCl adsorbed on H₂O ice

Coverage / (molecule cm ⁻²)	Temperature / K	Bibliographic Reference
5.0 10 ¹⁵	200	Hanson, D. R., Mauersberger, K., HCl/H ₂ O Solid Phase Vapor Pressures and HCl Solubility in Ice, <i>J. Phys. Chem.</i> , 94, 4700–4705, 1990
1.0 10 ¹⁵	200	Abbatt, J.P.D., Beyer, K. D., Fucaloro, A. F., McMahon, J. R., Wooldridge, P. J., Zhang, R., Molina, M. J., Interaction of HCl vapor with water ice: implications for the stratosphere, <i>J. Geophys. Res.</i> , 97, 15819–15826, 1992
(2.0 – 3.0) 10 ¹⁴	191	Hanson, D., Ravishankara, A.R., Investigation of the Reactive and Nonreactive Processes Involving ClONO ₂ and HCl on Water and Nitric Acid Doped Ice, <i>J. Phys. Chem.</i> 96, 2682-2691, 1992
1.15 10 ¹⁵	183	Foster, K. L., Tolbert, M. A., George, S. M., Interaction of HCl with Ice: Investigation of the Predicted Trihydrate, Hexahydrate, and Monolayer Regimes, <i>J. Phys. Chem. A</i> , 101, 4979–4986, 1997
2.5 10 ¹⁴	208	Interaction of HNO ₃ with water-ice surface at temperatures of the free troposphere, Abbatt, J.P.D., <i>Geophys Res. Lett.</i> 24, 1479-1482, 1997
3.1 10 ¹⁴	185	Flückiger, B., Thielmann, A., Gutzwiller, L., Rossi, M. J., Real time kinetics and thermochemistry of the uptake of HCl, HBr and HI on water ice in the temperature range 190 to 210 K, <i>Ber. Bunsenges. Phys. Chem.</i> , 102, 915–928, 1998
(1.1 ± 0.6) 10 ¹⁴	201	Lee, S.-H., Leard, D. C., Zhang, R., Molina, L. T., Molina, M. J., The HCl + ClONO ₂ reaction on various water ice surfaces, <i>Chem. Phys. Lett.</i> 315, 7–11, 1999
(2.0 ± 0.7) 10 ¹⁴	2001	Hynes, R. G., Mössinger, J. C., Cox, R. A.: The interaction of HCl with water-ice at tropospheric temperatures, <i>Geophys. Res. Lett.</i> 28, 2827–2830, 2001
1.7 10 ¹⁴ 1.3 10 ¹⁴ 6.7 10 ¹³	190 200 210	Flückiger, B., Rossi, M.J., Common Precursor Mechanism for the Heterogeneous Reaction of D ₂ O, HCl, HBr, and HOBr with Water Ice in the Range 170-230 K: Mass Accommodation Coefficients on Ice, <i>J. Phys. Chem. A</i> 197, 4103-4115, 2003
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