

23 reviewers' comments followed by our responses.

24 **Specific Comments:**

25 1. Line 93: “mixing” -> “mixed”

26 **Response:** Agreed and corrected in the revised manuscript (Line 94 in the revised
27 manuscript; Line 143 in this author's response).

28 2. Lines 102-103: can “under different air quality conditions” be changed to “under
29 pristine and polluted conditions”?

30 **Response:** Agreed and corrected in the revised manuscript (Line 104; Line 153).

31 3. Lines 108-109: “the radiative absorption enhancement by the aerosol mixtures in
32 East Asia has not been assessed.” is not accurate since there is a lot of work involved
33 in the radiative absorption enhancement, e.g., Cui et al. 2016 (doi:
34 10.1016/j.scitotenv.2016.02.026).

35 **Response:** According to the reviewer's comment and our research, we replaced “the
36 radiative absorption enhancement by the aerosol mixtures in East Asia has not been
37 assessed.” with a more proper description “further studies are urgently demanded to
38 better understand the key role that the East Asian aerosol mixtures play in the
39 formation mechanism of regional air pollution.” (Lines 109-111; Lines 158-160). We
40 cited Cui et al. 2016 (doi: 10.1016/j.scitotenv.2016.02.026) in the revised manuscript
41 (Lines 65-66 and Lines 426-428; Lines 65-66 and Lines 475-477).

42 4. Line 240: Fig. 4a -> Fig.3a

43 **Response:** Agreed and corrected in the revised manuscript (Line 242; Line 291).

44

45 **Additional changes**

46 The funding of Pengfei Tian was included in the revised manuscript (**Lines 394-395 in**
47 **the revised manuscript**; Lines 443-444 in this author's response), which was missing
48 in the original submitted manuscript.

49

50 **Radiative absorption enhancement of dust mixed with anthropogenic**
51 **pollution over East Asia**

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68

69 **ABSTRACT:** The particle mixing state plays a significant yet poorly quantified role
70 in aerosol radiative forcing, especially for the mixing of dust (mineral absorbing) and
71 anthropogenic pollution (black carbon absorbing) over East Asia. We have
72 investigated the absorption enhancement of mixed-type aerosols over East Asia by
73 using the Aerosol Robotic Network observations and radiative transfer model
74 calculations. The mixed-type aerosols exhibit significantly enhanced absorbing ability
75 than the corresponding unmixed dust and anthropogenic aerosols, as revealed in the
76 spectral behavior of absorbing aerosol optical depth, single scattering albedo, and
77 imaginary refractive index. The aerosol radiative efficiencies for the dust, mixed-type,
78 and anthropogenic aerosols are -101.0 , -112.9 and $-98.3 \text{ Wm}^{-2} \tau^{-1}$ at the bottom of the
79 atmosphere (BOA), -42.3 , -22.5 and $-39.8 \text{ Wm}^{-2} \tau^{-1}$ at the top of the atmosphere
80 (TOA), and 58.7 , 90.3 and $58.5 \text{ Wm}^{-2} \tau^{-1}$ in the atmosphere (ATM), respectively. The
81 BOA cooling and ATM heating efficiencies of the mixed-type aerosols are
82 significantly higher than those of the unmixed aerosol types over the East Asia region,
83 resulting in atmospheric stabilization. In addition, the mixed-type aerosols correspond
84 to a lower TOA cooling efficiency, indicating that the cooling effect by the
85 corresponding individual aerosol components is partially counteracted. We conclude
86 that the interaction between dust and anthropogenic pollution not only represents a
87 viable aerosol formation pathway but also results in unfavorable dispersion conditions,
88 both exacerbating the regional air pollution in East Asia. Our results highlight the
89 necessity to accurately account for the mixing state of aerosols in atmospheric models
90 over East Asia, in order to better understand the formation mechanism for regional air
91 pollution and to assess its impacts on human health, weather, and climate.

92

93 **1 Introduction**

94 Atmospheric aerosols or particulate matter (PM) profoundly affect the energy
95 budget of the earth-atmosphere system directly by interfering with the radiative
96 transfer and indirectly by modifying cloud formation (Twomey, 1977; Charlson and
97 Schwartz, 1992; Fan et al., 2007; Wang et al., 2011). However, the assessment of the
98 aerosol radiative effects is limited because of the inherent difficulties associated with
99 observations and model simulations (Stevens and Bony, 2013; Kok et al., 2017). In
100 particular, an accurate quantification of the mixing state of absorbing aerosols poses a
101 great challenge in the estimation of the aerosol direct radiative forcing (Haywood and
102 Boucher, 2000; He et al., 2015). Currently, the assessment of the aerosol direct and
103 indirect radiative forcing represents large uncertainty in the prediction of future
104 climate by anthropogenic activities (IPCC, 2007; IPCC, 2013). Moreover, aerosol
105 mixing state significantly affects atmospheric dynamics (Ramanathan and Carmichael,
106 2008).

107 Atmospheric aerosols are typically internally and/or externally mixed during
108 their lifetimes (Jacobson, 2001; Zhang and Zhang, 2005; Zhang et al., 2008; Khalizov
109 et al., 2009a; Pagels et al., 2009; Taylor et al., 2015). The East Asian region is
110 experiencing persistent heavy air pollution conditions in the present day (Guo et al.,
111 2014; Zhang et al., 2015; Wang et al., 2016). Black carbon (BC) is one of the major
112 anthropogenic pollutants in this region that exerts significant environmental and
113 climatic effects because of the strong absorption of solar radiation (Wang et al., 2012;
114 Peng et al., 2016), which is further enhanced by the “Lensing effect” (e.g., Jacobson,
115 2001; Cui et al., 2016). East Asia is also the largest dust source region second to the
116 Saharan Desert (Huang et al., 2014; Huang et al., 2015; Tian et al., 2015). As a result,
117 coarse mode dust particles are frequently mixed with anthropogenic pollution along

118 their transport pathway in East Asia (Noh et al., 2012; Logan et al., 2013; Guo et al.,
119 2017; Hara et al., 2017). The potential anthropogenic influence on dust has been
120 investigated close to the dust source regions (Huang et al., 2010; Bi et al., 2017).
121 Lower single scattering albedo (SSA) of mixed dust plumes has been assessed in
122 previous studies (Kim et al., 2005; Khatri et al., 2014). Li et al. (2015) have studied
123 the SSA spectral curvature of the East Asian aerosol mixtures using Aerosol Robotic
124 Network (AERONET) products and model simulations. The observations from a
125 sun-sky radiometer and a lidar have been applied to identify the presence of Asian
126 dust in mixed aerosol plumes at several East Asian monitoring sites (Noh et al., 2017).
127 Those earlier studies on optical properties and radiative effects of the East Asian dust
128 and anthropogenic aerosol mixtures have promised reduction of the uncertainties in
129 estimating the aerosol radiative effects.

130 Internal mixing of coarse mode dust with fine mode anthropogenic aerosols has
131 been suggested by observations in the Asian Aerosol Characterization Experiment
132 (ACE-Asia) (Seinfeld et al., 2004) and recent studies (Sugimoto et al., 2015; Wang et
133 al., 2017), although the mechanism leading to the mixing has yet to be elucidated.
134 Internal mixing of dust with anthropogenic pollution likely occurs via condensation of
135 low-volatility organic and inorganic compounds, particle-phase reactions, and
136 coagulation with other aerosol types (Zhang et al., 1996; Zhao et al., 2006; Qiu et al.,
137 2011). In addition, dust particles provide reactive surfaces for catalytic conversion of
138 sulfur dioxide to sulfate, which has been suggested as a key mechanism for severe
139 haze formation in China (Zhang et al., 2015; Li et al., 2017a). Atmospheric
140 measurements using electron microscopy have identified BC and certain soluble
141 aerosols on the surface of dust particles (Tobo et al., 2010; Ma et al., 2012; Li et al.,
142 2014). Pan et al. (2017) have studied the morphology change of East Asian dust

143 **mixed** with anthropogenic aerosols and showed the possibility evidence for the
144 occurrence of aqueous-phase reactions.

145 The aerosol mixing state significantly affects the radiative effects (Jacobson,
146 2001; Khalizov et al., 2009b; Xue et al., 2009). Several previous studies have shown
147 that the amount of solar radiation reaching the Earth surface through mixtures of
148 mineral dust and other absorbing aerosols is considerably reduced compared to that
149 through dust-only aerosols (Derimian et al. 2008; Obregón et al. 2015). Researchers
150 have reported that the radiative efficiency of non-dust aerosols is higher than that of
151 dust aerosols at two urban Asian cities of Gwangju and Beijing (Noh et al., 2012; Yu
152 et al., 2016). Maximum radiative efficiency under unpolluted conditions has been
153 found by comparing aerosol radiative effects under **pristine and polluted** conditions
154 (Chen et al., 2016). Aerosol radiative efficiency has been found to be strongly
155 influenced by aerosol absorbing ability (e.g., SSA) and size of fine mode particles in
156 Central China (Zhang et al., 2017).

157 Although the mixing state of dust and anthropogenic aerosols considerably
158 affects aerosol radiative effects, **further studies are urgently demanded to better**
159 **understand the key role that the East Asian aerosol mixtures play in the formation**
160 **mechanism of regional air pollution.** In this present work, we have extensively
161 investigated the radiative absorption enhancement by the East Asian aerosol mixtures
162 on the basis of long-term AERONET observations and SBDART model simulations.
163 We classified the dust, mixed-type, and anthropogenic aerosols according to the
164 aerosol SSA spectral behavior and Ångström exponent parameter. The optical and
165 microphysical properties of the various aerosol types were analyzed, with emphasis
166 on the absorption enhancement by mixed-type aerosols. The mechanism leading to the
167 radiative absorption enhancement by the mixed-type aerosols and their impacts on

168 regional air pollution and climate have been discussed. Our results suggest that the
169 East Asian aerosol mixtures result in a more stable atmosphere that is unfavorable for
170 diffusion and dispersion of the atmospheric pollutants.

171 **2 Data and methodology**

172 **2.1 AERONET data**

173 The aerosol optical and microphysical data used in this research were originally
174 from the Aerosol Robotic Network (AERONET) (Holben et al., 1998). To ensure the
175 data quality, we only analyzed the cloud-screened, quality-assured Level 2.0 inversion
176 data. We further constrained the data with solar zenith angle between 50° and 80° to
177 avoid possible inversion errors and surface albedo smaller than 0.5 to exclude
178 seasonal snow-covered surfaces. The AERONET spectral products are available at the
179 wavelengths of 440, 675, 870, and 1020 nm, respectively. Previously, Dubovik et al.
180 (2000) have evaluated the uncertainty of the AERONET products.

181 All available worldwide measurement sites from the Aerosol Robotic Network
182 (AERONET) program with a sample number of greater than 100 were included in the
183 present work. Considering the regional representativeness and data availability, 11
184 sites (Table S1 and Fig. S1) were selected to represent the various types of the East
185 Asian aerosol mixtures. For example, the sample numbers of Beijing and Xianghe
186 were randomly reduced to a one-fourth of the total numbers because these two sites
187 are close to each other and the total sample numbers of the two sites are significantly
188 larger than the others.

189 **2.2 Radiative forcing and efficiency calculations**

190 The aerosol direct radiative forcing (ΔF) is defined as follows:

$$191 \quad \Delta F = (F^\downarrow - F^\uparrow) - (F_0^\downarrow - F_0^\uparrow) \quad (1)$$

192 where F and F_0 are radiative fluxes under the aerosol-free and aerosol-laden

193 conditions respectively, the upward and downward arrows denote the directions of the
194 radiative fluxes. The AERONET products include the aerosol radiative forcing and
195 efficiency, in addition to the optical and microphysical parameters. However, the
196 radiative forcing in the AERONET products was not exactly calculated using
197 Equation (1). Instead, we calculated the aerosol direct radiative forcing and efficiency
198 by using the widely adopted Santa Barbara DISORT Atmospheric Radiative Transfer
199 (SBDART) model (Ricchiuzzi et al., 1998). The aerosol radiative efficiency (ΔF^{eff}) is
200 defined as the aerosol direct radiative forcing per unit aerosol optical depth (AOD):

$$201 \quad \Delta F^{eff} = \Delta F / AOD_{0.55} \quad (2)$$

202 where $AOD_{0.55}$ is the 550 nm AOD. However, the aerosol direct radiative forcing is
203 not linearly dependent on AOD (e.g., Wu et al., 2015). To exclude possible errors
204 from aerosol loading in calculating the aerosol radiative efficiency in Equation (2), the
205 550 nm AOD was set to unity ($AOD_{0.55}=1.0$) in SBDART model calculations. The data
206 processing procedure is presented in Fig. 1 and a detailed discussion of aerosol
207 radiative forcing and efficiency calculation has been provided elsewhere (Tian et al.,
208 2018).

209 **2.3 Aerosol classification**

210 The SSA spectral behavior and the Ångström exponent parameter were applied
211 to identify aerosol mixtures (Fig. 1). The main absorbing component of the
212 anthropogenic pollutants is BC, which exhibits strong absorption throughout a broad
213 wavelength range (Bond and Bergstrom, 2006). Dust aerosols are also an absorbing
214 medium because of the iron-bearing minerals such as hematite, goethite and magnetite,
215 which enable the dust aerosols to absorb strongly in the UV and short visible
216 wavelengths and weakly in the near infrared (Schuster et al., 2016). Thus, the dust
217 aerosols exhibit a monotonically increasing SSA trend with increasing wavelength,

218 while the anthropogenic pollutants (the absorption of which is mainly contributed by
219 BC) show a monotonically decreasing SSA spectra (Bergstrom et al., 2007; Giles et
220 al., 2012; Tian et al., 2017). The absorption of aerosol mixtures is contributed by both
221 dust and BC, leading to a non-monotonic SSA spectra (Khatri et al., 2014). The
222 characteristic non-monotonic SSA spectral behavior of mixed-type aerosols provides
223 a useful approach to identify the mixed-type aerosols. SSA curvature of greater than
224 0.1, which has been suggested for “moderate mixing” (Li et al., 2015), was applied
225 for the mixed-type aerosol. We employed an additional criterion, Ångström exponent,
226 which is related to the aerosol size and the widely used in aerosol classification (Eck
227 et al., 2010; Derimian et al., 2016), to further constrain the aerosol classification.
228 Results show that Ångström exponent values are smaller for coarse mode dust
229 aerosols (lower than 0.7 in the present study), larger for fine mode anthropogenic
230 aerosols (greater than 1.4 in this research), and range from 0.8 and 1.3 for the
231 mixed-type aerosols, respectively.

232 The imaginary refractive index at 440 nm (k_{440} , in the visible bands) versus the
233 average of the imaginary refractive indices at 675, 870, and 1020 nm (k_{mir} , in the red
234 and near-infrared bands) is shown in Fig. 2. Anthropogenic aerosols with the
235 absorption mainly contributed by BC particles exhibit a flat imaginary refractive
236 index (Bond and Bergstrom, 2006), while dust with the components such as hematite
237 absorbs strongly in UV but weakly in longer wavelengths (Hsu and Matijević, 1985;
238 Schuster et al., 2016). Hence, the data of anthropogenic aerosols are scattered along
239 the 1:1 line (Fig. 2). Dust aerosols show stronger absorption at the visible wavelength
240 than at the red and near-infrared wavelengths, so the data are scattered above the 1:1
241 line. The data for most of the mixed-type aerosols lie above the 1:1 line and on the
242 right side of the $k_{mir} = 0.0042$ threshold, which is suggested by Schuster et al. (2016)

243 to separate dust ($k_{mir} < 0.0042$) and biomass burning ($k_{mir} > 0.0042$) aerosols. The
244 mixed-type aerosols show the strongest absorption and most of the sites with a
245 mixed-type aerosol sample number of 50 and higher are located in East Asia. Hence,
246 our method quantitatively separated various aerosol types.

247 **3 Spectral behavior of the East Asian aerosol mixtures**

248 Extensive investigations of the spectral optical properties were carried out to
249 discuss the enhanced absorbing ability of the East Asian aerosol mixtures. Dust
250 aerosols show the highest spectral AOD and the anthropogenic aerosols exhibit the
251 lowest spectral AOD (Fig. 3a). The spectral dependence of dust aerosols is nearly
252 invariant across the wavelength spectrum while the dependence of anthropogenic
253 aerosols is relatively high, which is relevant to aerosol size and the Ångström
254 parameter (Ångström, 1929). The mixed-type aerosols have smaller AOD than dust
255 aerosols, but higher absorption aerosol optical depth (AAOD) throughout the
256 wavelength band of 440 to 1020 nm (Fig. 3b). Dust aerosols exhibit higher AAOD at
257 the 440 nm wavelength than that of anthropogenic aerosols.

258 As expected by using our SSA spectral classification method, dust aerosols show
259 a monotonic increasing SSA trend with increasing wavelength, while anthropogenic
260 aerosols exhibit an opposite trend. In contrast, the SSA of the mixed-type aerosols
261 peaks at the wavelength of 675 nm (Fig. 3c). Interestingly, the mixed-type aerosols
262 exhibit the lowest SSA value that cannot be predicted by our classification method
263 because the classification method considers the spectral trend rather than the value of
264 SSA. This indicates enhanced absorption for the mixed-type aerosols. The spectral
265 average SSA of the dust, mixed-type, and anthropogenic aerosols are 0.94, 0.89, and
266 0.93, respectively. Hence, internal mixing of the East Asian aerosol mixtures yields
267 the lowest SSA that is distinct from the corresponding individual aerosol types.

268 The dust aerosols exhibit the highest value for the real part of the complex
269 refractive index, while the anthropogenic aerosols show the lowest value (Fig. 3e).
270 Interestingly, the real refractive index of the mixed-type aerosols is close to that of
271 dust aerosols, indicating high scattering of the mixed-type aerosols. The spectral
272 imaginary refractive index for anthropogenic aerosols is nearly constant (Fig. 3f),
273 which is characteristic of this aerosol type. However, the imaginary refractive index
274 of anthropogenic aerosols is much lower than that of BC aerosols (approximately 0.6
275 in Bond and Bergstrom (2006)), because the majority of anthropogenic aerosol
276 components are non-absorbing aerosols such as sulfate and nitrate salts. The spectral
277 imaginary refractive index of dust aerosols is similar to the AERONET dust
278 climatology over Africa and the Middle East (Schuster et al., 2016), with stronger
279 absorption at the visible wavelength (440 nm). The mixed-type aerosols exhibit the
280 highest imaginary refractive index, especially at the visible wavelength, where the
281 imaginary refractive index of the mixed-type aerosols (0.0159) is more than twice that
282 of anthropogenic aerosols (0.0078).

283 Hence, the absorbing ability of the East Asian aerosols is significantly enhanced
284 due to the mixing process, in light of the lowest SSA and the highest AOD and
285 imaginary refractive index for the mixed-type aerosols.

286 **4 Enhanced radiative absorption by East Asian aerosol mixtures**

287 To investigate the radiative effects caused by the enhanced absorbing ability of
288 the East Asian aerosol mixtures, we calculated the average aerosol direct radiative
289 efficiency at the bottom of the atmosphere (BOA), at the top of the atmosphere (TOA),
290 and in the atmosphere (ATM) (Fig. 4), respectively. The mixed-type aerosols exhibit
291 lower spectral average AOD (0.48) than dust aerosols (0.64) (Fig. 3a), but show
292 comparable radiative forcing relative to dust aerosols at BOA (-72.7 and -73.6 Wm⁻²

293 for the mixed-type and dust aerosols, respectively) (Fig. 4a). This feature is explained
294 by higher BOA cooling efficiency of the mixed-type aerosols ($-112.9 \text{ Wm}^{-2}\tau^{-1}$) than
295 dust ($-101.0 \text{ Wm}^{-2}\tau^{-1}$) (Fig. 4b). The radiative absorption enhancement is evident for
296 the TOA and ATM forcing: the mixed-type aerosols exhibit the highest ATM radiative
297 forcing (55.0 Wm^{-2}) and the lowest absolute TOA forcing (-17.8 Wm^{-2}). For
298 comparison, we calculated the aerosol radiative efficiency of various aerosol types to
299 rule out the effect of aerosol loading. The mixed-type aerosols exhibit the highest
300 BOA cooling efficiency, the highest ATM heating efficiency ($90.3 \text{ Wm}^{-2}\tau^{-1}$), and the
301 lowest TOA cooling efficiency ($-22.5 \text{ Wm}^{-2}\tau^{-1}$).

302 The average BOA radiative efficiency of dust aerosols ($-101.0 \text{ Wm}^{-2}\tau^{-1}$) in the
303 present study (Fig. 4b) falls in the range of -96.1 to $-127.0 \text{ Wm}^{-2}\tau^{-1}$ by Yu et al. (2016),
304 but lower than the result of $-124.6 \pm 12.2 \text{ Wm}^{-2}\tau^{-1}$ by Noh et al. (2012). Note that
305 results by Noh et al. (2012) and Yu et al. (2016) are likely biased due to a non-linear
306 dependence between aerosol direct radiative forcing and AOD, which is avoided in
307 the present study. Our previous work (Tian et al., 2018), which used the same
308 radiative efficiency calculation method but a different aerosol classification approach
309 from the present study, obtained a similar BOA radiative efficiency of $-102.3 \text{ Wm}^{-2}\tau^{-1}$
310 at the Semi-Arid Climate and Environment Observatory of Lanzhou University
311 (SACOL), Northwest China.

312 The spatial distributions of the aerosol radiative efficiency for BOA, ATM, and
313 TOA are presented in Figs. 5-7, respectively. Only those sites with a sample number
314 of greater than 50 were averaged and included in the figures. Despite the fact that
315 AERONET sites are unavailable in some remote areas, the worldwide aerosol
316 distributions are well captured by the AERONET observations. The mixed-type
317 aerosols are distributed in East Asia, India and around the Saharan Desert regions.

318 The mixed-type aerosols exhibit a higher BOA radiative cooling efficiency than
319 that of dust and anthropogenic aerosols over the East Asia region (Fig. 5). The BOA
320 radiative cooling efficiency over India is also high, but the difference of various
321 aerosol types is small. Biomass burning aerosols over Africa exhibit the highest BOA
322 cooling efficiency in the globe, which may be explained by the different combustion
323 compositions and processes as described in Eck et al. (2010) and Garc ía et al. (2012).
324 Anthropogenic aerosols over North America and dust aerosols around the Sahara
325 Desert show a relatively lower BOA cooling efficiency.

326 The mixed-type aerosols also exhibit a higher ATM radiative heating efficiency
327 than dust and anthropogenic aerosols over East Asia (Fig. 6). The ATM radiative
328 heating efficiency over India is high for all aerosol types. The ATM radiative heating
329 efficiency is high over South Africa, where biomass burning aerosols dominate. On
330 the other hand, the ATM radiative heating efficiency over North America and around
331 the Sahara Desert regions is relatively smaller.

332 The enhanced BOA cooling and ATM heating efficiencies reveal that the
333 mixed-type aerosols exhibit higher BOA cooling and ATM heating effects than those
334 of the unmixed dust and anthropogenic aerosols with the same aerosol loading. The
335 enhanced BOA cooling and ATM heating effects lead to a cooler surface and warmer
336 atmosphere and restrain the development of the planetary boundary layer, resulting in
337 a more stable atmosphere that is unfavorable for dispersion of atmospheric gaseous
338 and PM pollutants. Noting that mixed-type aerosols occur frequently in East Asia and
339 their occurrence can reach as high as fifty percent over some locations (Li et al.,
340 2015). Hence, the mixed-type aerosols likely play a significant role in enhancing air
341 pollution over East Asia. In addition, the mixed-type aerosols show lower TOA
342 radiative cooling efficiency than dust and anthropogenic aerosols over the East Asia

343 region (Fig. S1). The reduced TOA cooling efficiency indicates that the East Asian
344 aerosol mixtures partially counteract the cooling effect of the Earth-atmosphere
345 system by the corresponding individual components.

346 **5 Discussions**

347 The aerosol direct radiative efficiency strongly depends on solar zenith angle
348 (e.g., Derimian et al., 2016). To investigate the influence of solar zenith angle on the
349 result of the present study, the aerosol radiative efficiency as a function of solar zenith
350 angle were calculated for various aerosol types (Fig. 7). Note that the AERONET data
351 used in the present study is only available between 50 ° to 80 ° solar zenith angles. The
352 BOA radiative cooling and ATM heating efficiencies (absolute value of radiative
353 efficiency) decrease with increasing solar zenith angle (decreasing cosine of solar
354 zenith angle), while the TOA cooling efficiency increase with increasing solar zenith
355 angle. The mixed-type aerosols exhibit higher BOA cooling efficiency, higher ATM
356 heating efficiency, and lower TOA cooling efficiency than those of unmixed dust and
357 anthropogenic aerosols. The dust aerosols exhibit both higher BOA and TOA cooling
358 efficiency than the anthropogenic aerosols, leading to small difference in the ATM
359 heating efficiency between the two aerosol types.

360 Aerosol absorption has been suggested as a key factor that determines the aerosol
361 radiative effects (Li et al., 2010). The aerosol radiative efficiencies as a function of
362 SSA and imaginary refractive index for various aerosol types were calculated to
363 investigate the effect of aerosol absorbing on the aerosols radiative efficiency (Figs. 8
364 and 9). The BOA cooling efficiency and ATM heating efficiency increase with
365 increasing absorption (i.e., decreasing SSA and increasing imaginary refractive index)
366 and the TOA cooling efficiency decrease with increasing absorption. However, the
367 dependences between radiative efficiency and SSA are stronger than those between

368 radiative efficiency and the imaginary refractive index for BOA, TOA, and ATM. The
369 dependences between radiative efficiency and SSA are approximately linear, but the
370 dependences between radiative efficiency and imaginary refractive index become less
371 apparent with increasing imaginary refractive index. The strong dependence between
372 aerosol radiative efficiency and SSA has also been shown over Central China and
373 desert and semi-desert regions of northwestern China (Xin et al., 2016; Zhang et al.,
374 2017).

375 We also examined the effects of the fraction of fine and coarse mode on aerosol
376 radiative efficiency (Fig. 10). The BOA cooling and ATM heating efficiencies (TOA
377 cooling efficiency) initially increase (decreases) with increasing fine mode fraction
378 (FMF), when FMF is lower than 0.3 and coarse mode dust aerosols dominate. Then
379 the BOA cooling and ATM heating efficiencies (TOA cooling efficiency) reach a peak
380 (bottom) in the FMF range of 0.3 to 0.5. Finally the BOA cooling and ATM heating
381 efficiency (TOA cooling efficiency) begin to decrease (increase) when FMF is greater
382 than 0.5, where fine mode anthropogenic aerosols become dominate. Overall, the
383 moderate mixing of dust with fine mode anthropogenic pollutants, which is classified
384 as mixed-type aerosols in the present study, is responsible for the enhanced radiative
385 absorption. A previous study (Li et al., 2015) has revealed that moderate mixing of
386 East Asian dust with fine mode pollutants is responsible for high SSA spectral
387 curvature, which suggests well-mixed aerosol mixtures.

388 **6 Conclusions**

389 The mixing state of atmospheric aerosols plays a significant yet poorly
390 quantified role in determining the aerosol optical properties and radiative effects. In
391 the East Asia region, coarse mode dust and fine mode anthropogenic pollution are
392 typically mixed externally and/or internally in the atmosphere. The mixing of dust

393 with anthropogenic aerosols exerts a significant influence on aerosol absorption and
394 radiative efficiency. We present an extensive investigation of the radiative effects of
395 the East Asian aerosol mixtures.

396 The mixed-type aerosols exhibit significantly higher AAOD, lower SSA, and
397 higher imaginary refractive index than those of unmixed dust and anthropogenic
398 aerosols, showing significantly enhanced absorption. The absorption enhancement is
399 most evident at wavelength 440 nm, where the imaginary refractive index of the
400 mixed-type aerosols (0.0159) is more than twice that of anthropogenic aerosols
401 (0.0078). The mixed-type aerosols also exhibit a unique non-monotonic SSA trend,
402 which provides a characteristic signature for identifying these aerosols.

403 The values of the aerosol radiative efficiencies for dust, mixed-type and
404 anthropogenic aerosols are -101.0, -112.9 and -98.3 $\text{Wm}^{-2}\tau^{-1}$ for BOA, -42.3, -22.5
405 and -39.8 $\text{Wm}^{-2}\tau^{-1}$ for TOA, and 58.7, 90.3 and 58.5 $\text{Wm}^{-2}\tau^{-1}$ for ATM, respectively.
406 The mixed-type aerosols exhibit significantly higher BOA radiative cooling efficiency
407 and ATM heating efficiency than those of dust and anthropogenic aerosols over East
408 Asia. These enhanced BOA cooling and ATM heating efficiencies reveal that the
409 mixed-type aerosols exhibit stronger BOA cooling and ATM heating effects than
410 those of unmixed dust and anthropogenic aerosols for a given aerosol loading,
411 resulting in a more stable atmosphere that is unfavorable for the diffusion and
412 dispersion of the gaseous and PM pollution. Hence, our results suggest that the
413 mixed-type aerosols likely play a significant role in enhancing the air pollution in East
414 Asia, because the mixing of dust and anthropogenic aerosols occurs frequently in this
415 region. In addition, the mixed-type aerosols show lower TOA cooling efficiency,
416 indicating that the mixed-type aerosols partially counteract the cooling effect of the
417 earth-atmosphere system by the corresponding individual components. Since dust

418 particles play a catalytic role in the conversion of sulfur dioxide to sulfate (Zhang et
419 al., 2015; Li et al., 2017a), the interaction between dust and anthropogenic aerosols
420 also provides a possible mechanism for the observed efficient internal mixing and
421 enhanced radiative absorption over East Asia.

422 Multiple factors have been suggested to be responsible for severe air pollution in
423 East Asia, including the interaction between BC aerosols and the atmospheric
424 boundary layer (Ding et al., 2016; Peng et al., 2016; Li et al., 2017b), rapid secondary
425 aerosol formation during severe haze events (Wang et al., 2016), weakening of the
426 East Asian monsoon circulation (Wu et al., 2016), and climate change (Cai et al,
427 2017). Our results indicate that interaction between dust and anthropogenic pollution
428 does not only represent a plausible pathway for PM formation and internal mixing but
429 also results in unfavorable dispersion conditions (i.e., increased atmospheric stability),
430 both exacerbating regional air pollution in East Asia. Clearly, future studies are
431 necessary to more accurately assess the mixing state of aerosols in atmospheric
432 models, in order to better understand the formation mechanism for air pollution over
433 East Asian and to assess its impacts on human health, weather, and climate (Zhang et
434 al., 2015).

435 **7 Data availability**

436 The original sun photometer data are available from the AERONET website
437 (<http://aeronet.gsfc.nasa.gov/>). The radiative flux data for the worldwide AERONET
438 sites calculated using the SBDART model and all data for the figures and table in the
439 present research are available from the authors upon request.

440 **Competing interests.** The authors declare that they have no conflict of interest.

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Figure Captions

719 **Figure 1.** Diagram of the radiative efficiency calculation and aerosol classification

720 **Figure 2.** The imaginary refractive index at 440 nm versus the average of the imaginary refractive
721 indices at 675, 870, and 1020 nm for dust, mixed-type, and anthropogenic aerosols over
722 worldwide AERONET sites. Only those sites with a sample number of 50 and higher were
723 averaged and shown in the figure.

724 **Figure 3.** Spectral behavior of (a) AOD, (b) AAOD, (c) SSA, (d) asymmetry factor, (e) real part of
725 the complex refractive index, and (f) imaginary part of the complex refractive index for dust,
726 mixed-type and anthropogenic aerosols averaged from East Asian sites.

727 **Figure 4.** (a) Radiative forcing and (b) radiative efficiency of the dust, mixed-type, and
728 anthropogenic aerosols averaged from East Asian sites.

729 **Figure 5.** Aerosol radiative efficiency at BOA: (a) dust aerosols, (b) anthropogenic aerosols, and
730 (c) mixed-type aerosols.

731 **Figure 6.** Aerosol radiative efficiency in ATM: (a) dust aerosols, (b) anthropogenic aerosols, and
732 (c) mixed-type aerosols.

733 **Figure 7.** Aerosol direct radiative efficiency as a function of the cosine of solar zenith angle for
734 the dust, mixed-type and anthropogenic aerosols: (a) BOA, (b) ATM, and (c) TOA.

735 **Figure 8.** Aerosol direct radiative efficiency as a function of SSA for the dust, mixed-type and
736 anthropogenic aerosols: (a) BOA, (b) ATM, and (c) TOA.

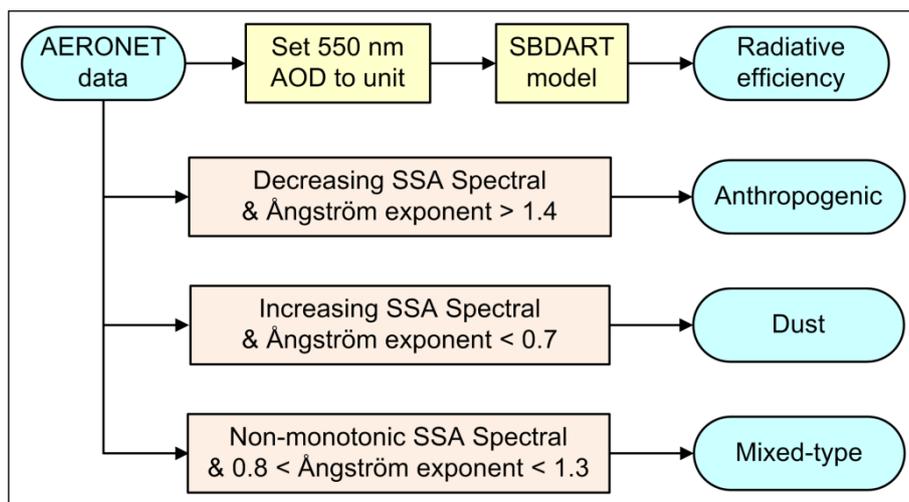
737 **Figure 9.** Aerosol direct radiative efficiency as a function of imaginary refractive index for the
738 dust, mixed-type and anthropogenic aerosols: (a) BOA, (b) ATM, and (c) TOA.

739 **Figure 10.** Aerosol direct radiative efficiency as a function of FMF in the East Asian region: (a)
740 BOA, (b) ATM, and (c) TOA. The average radiative efficiencies of the dust, mixed-type and
741 anthropogenic aerosols are also plotted.

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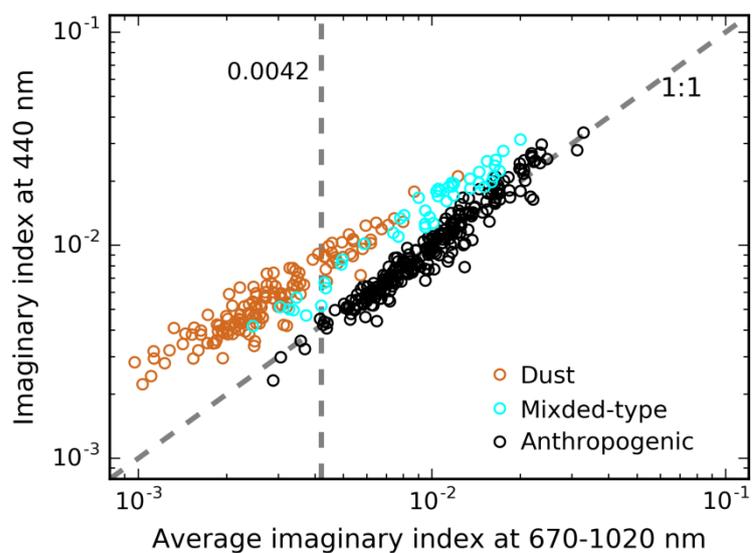
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Figure 1. Diagram of the radiative efficiency calculation and aerosol classification

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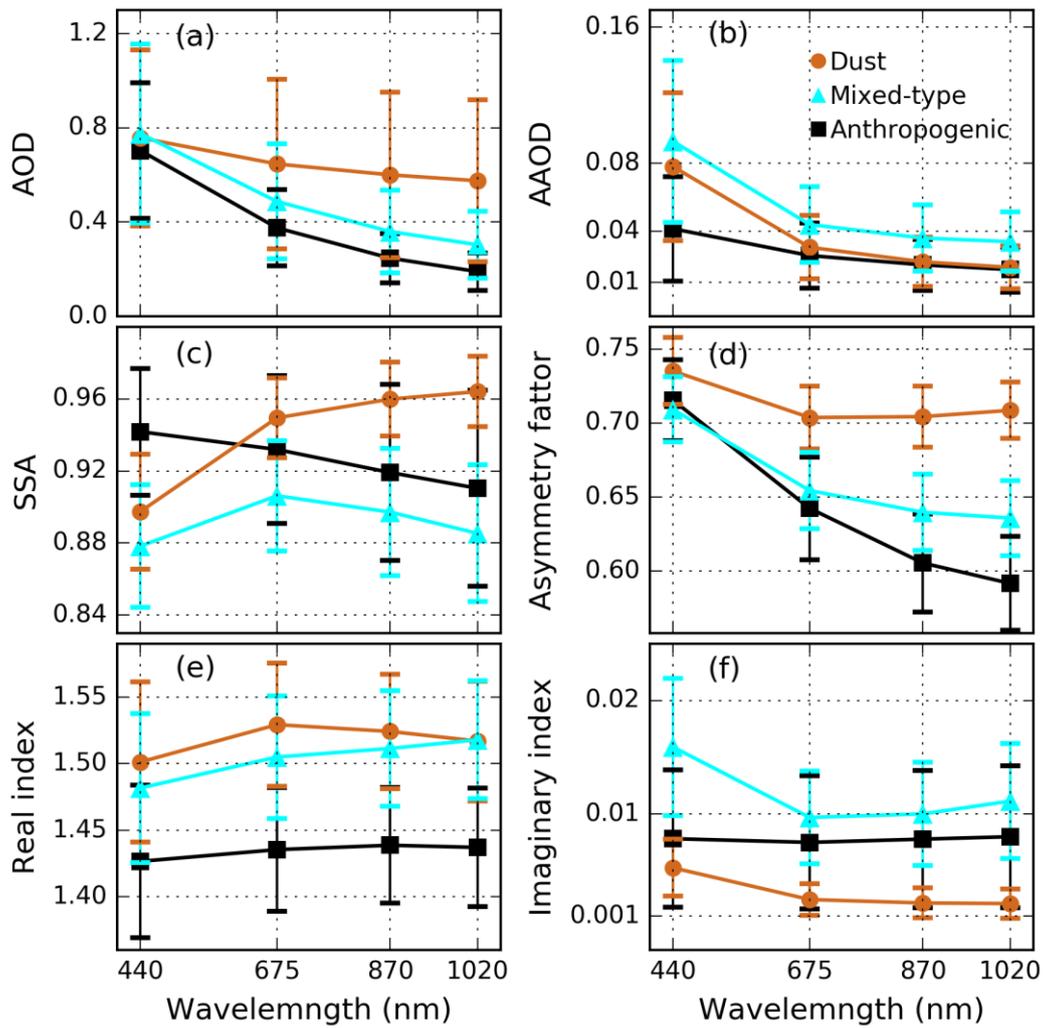
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750 **Figure 2.** The imaginary refractive index at 440 nm versus the average of the imaginary refractive
751 indices at 675, 870, and 1020 nm for dust, mixed-type, and anthropogenic aerosols over
752 worldwide AERONET sites. Only those sites with a sample number of 50 and higher were
753 averaged and shown in the figure.

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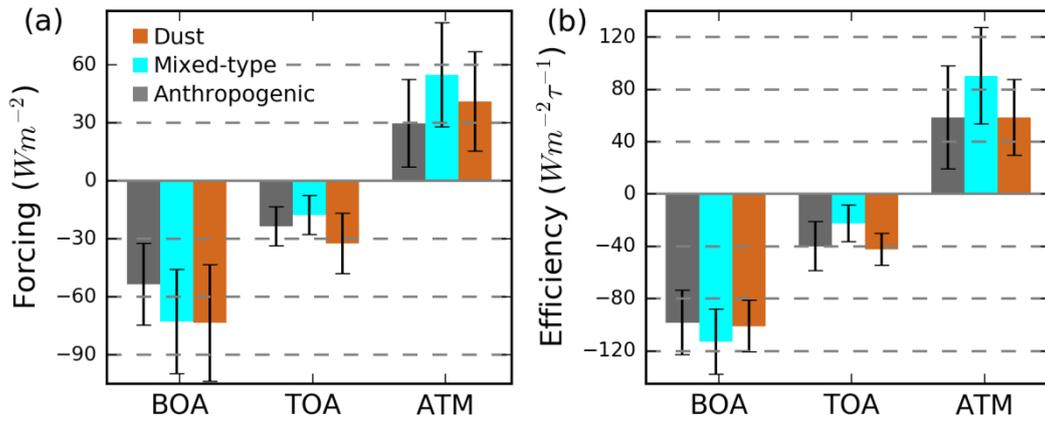


756

757 **Figure 3.** Spectral behavior of (a) AOD, (b) AAOD, (c) SSA, (d) asymmetry factor, (e) real part of
 758 the complex refractive index, and (f) imaginary part of the complex refractive index for dust,
 759 mixed-type and anthropogenic aerosols averaged from East Asian sites.

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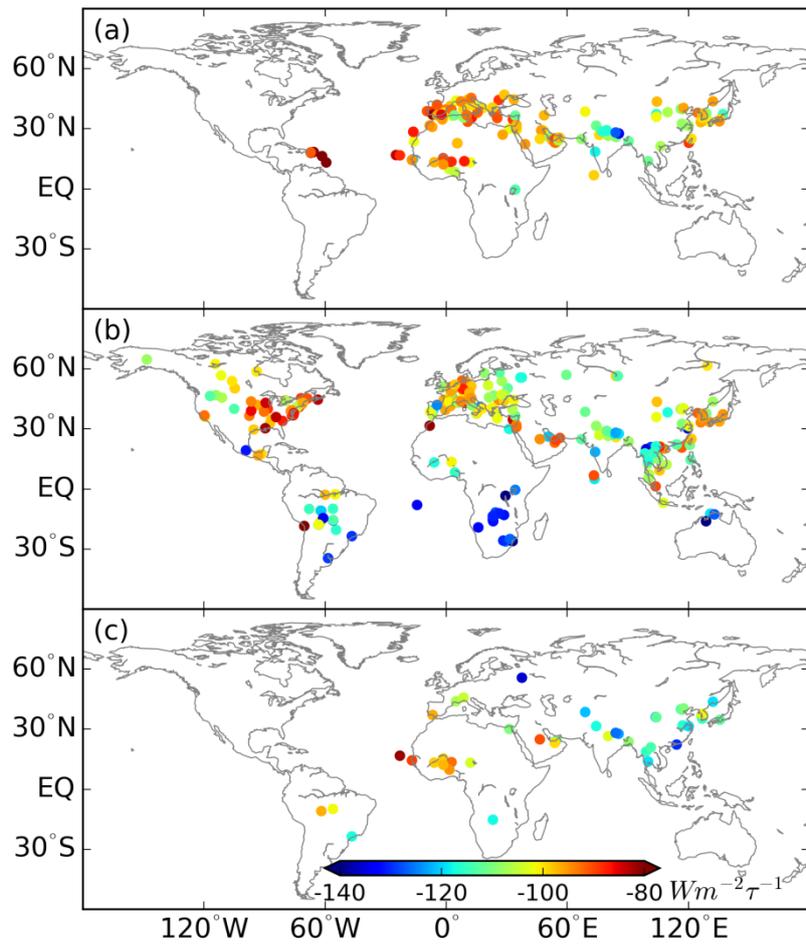
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762

763 **Figure 4.** (a) Radiative forcing and (b) radiative efficiency of the dust, mixed-type, and
764 anthropogenic aerosols averaged from East Asian sites.

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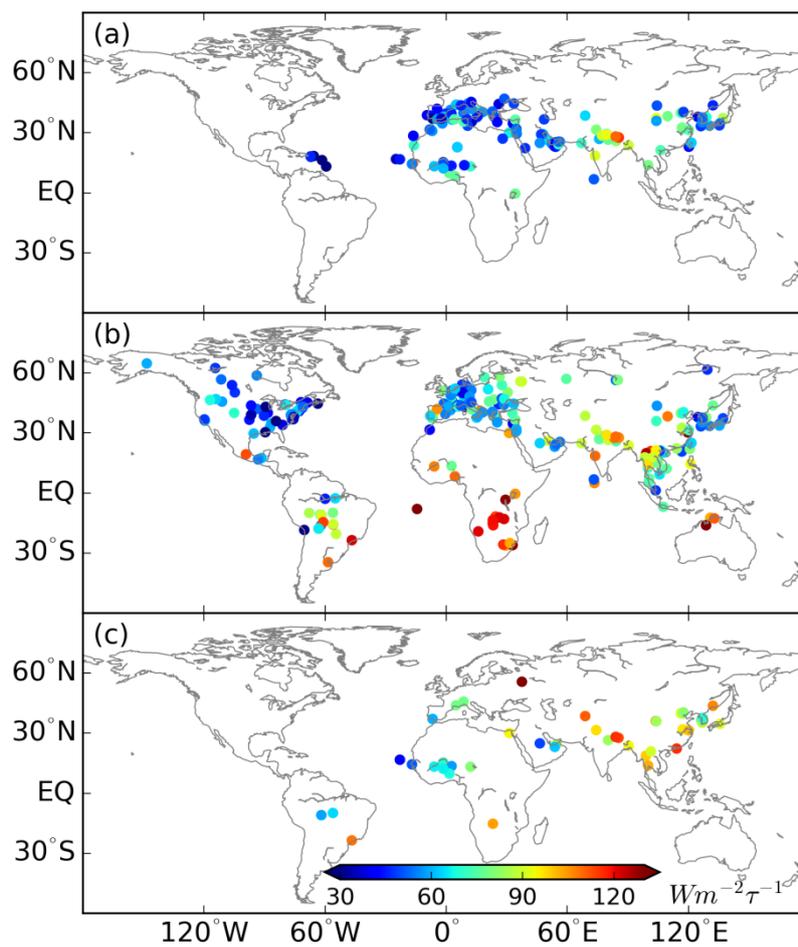
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768 **Figure 5.** Aerosol radiative efficiency at BOA: (a) dust aerosols, (b) anthropogenic aerosols, and

769 (c) mixed-type aerosols.

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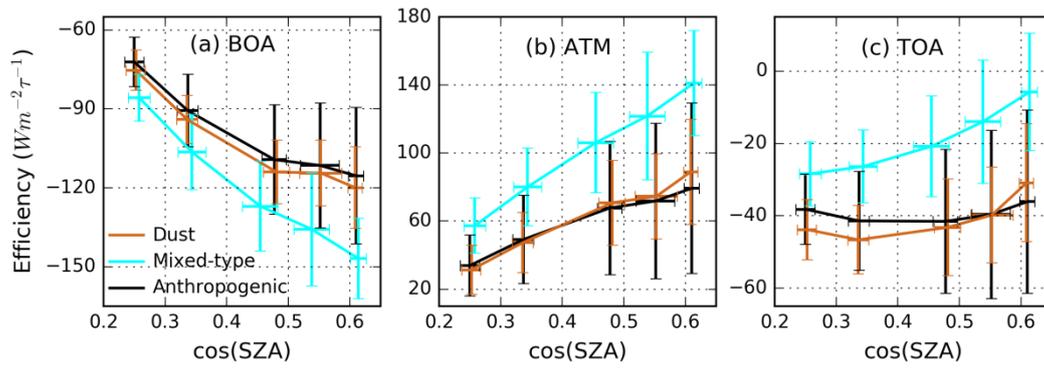
773

774 **Figure 6.** Aerosol radiative efficiency in ATM: (a) dust aerosols, (b) anthropogenic aerosols, and

775 (c) mixed-type aerosols.

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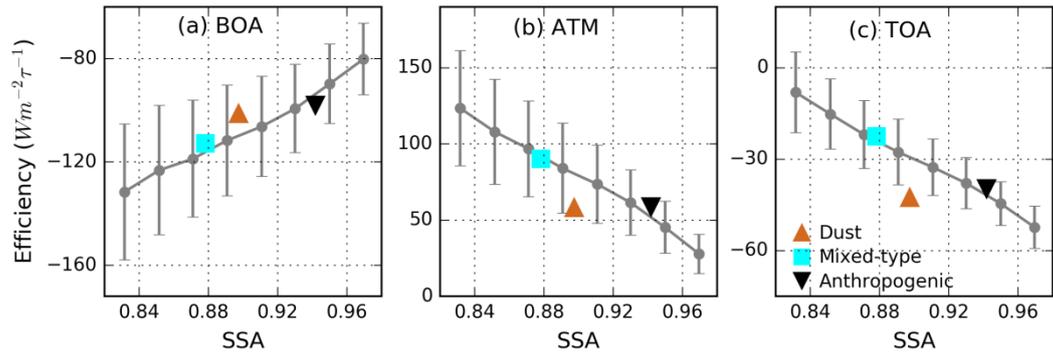


778

779 **Figure 7.** Aerosol direct radiative efficiency as a function of the cosine of solar zenith angle for
780 the dust, mixed-type and anthropogenic aerosols: (a) BOA, (b) ATM, and (c) TOA.

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Figure 8. Aerosol direct radiative efficiency as a function of SSA for the dust, mixed-type and

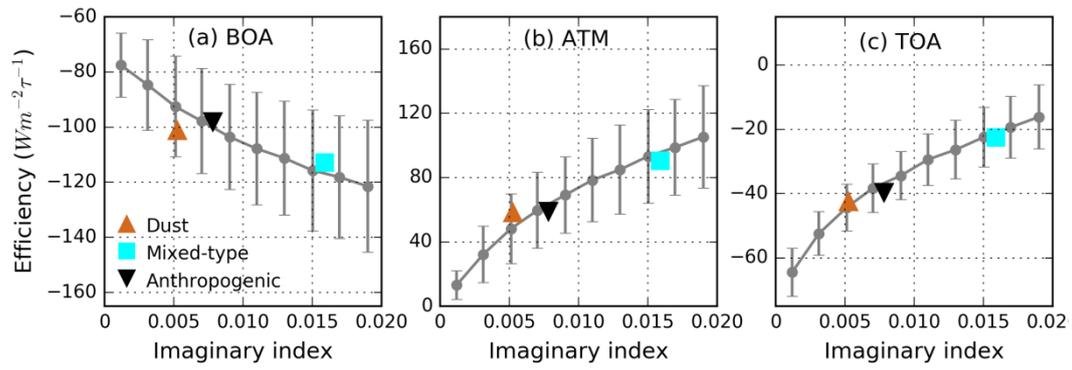
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anthropogenic aerosols: (a) BOA, (b) ATM, and (c) TOA.

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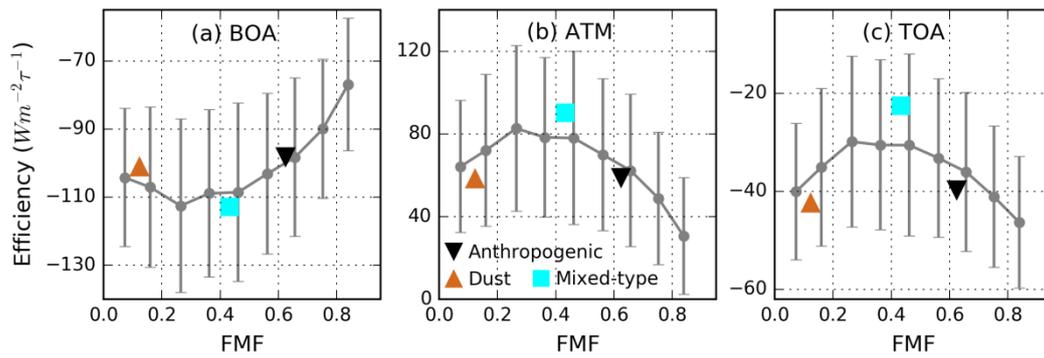


789

790 **Figure 9.** Aerosol direct radiative efficiency as a function of imaginary refractive index for the
791 dust, mixed-type and anthropogenic aerosols: (a) BOA, (b) ATM, and (c) TOA.

792

793



794

795 **Figure 10.** Aerosol direct radiative efficiency as a function of FMF in the East Asian region: (a)

796 BOA, (b) ATM, and (c) TOA. The average radiative efficiencies of the dust, mixed-type and

797 anthropogenic aerosols are also plotted.

798