

1 **Differentiating between particle formation and growth events in an urban**
2 **environment**

3
4
5 Buddhi Pushpawela, Rohan Jayaratne, and Lidia Morawska *

6
7
8 International Laboratory for Air Quality and Health

9 Queensland University of Technology

10 GPO Box 2434, Brisbane, QLD 4001, Australia

11
12 Revised and submitted to

13 Atmospheric Chemistry and Physics

14 June 2018

15
16
17 * Corresponding author contact details:

18 Tel: (617) 3138 2616; Fax: (617) 3138 9079

19 Email: lmorawska@qut.edu.au

Abstract

22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41

Small aerosols at a given location in the atmosphere often originate in-situ from new particle formation (NPF). However, they can also be produced and then transported from a distant location to the point of observation where they may continue to grow to larger sizes. This study was carried out in the subtropical urban environment of Brisbane, Australia, in order to assess the relative occurrence frequencies of NPF events and particle growth events with no NPF. We used a neutral cluster and air ion spectrometer (NAIS) to monitor particles and ions in the size range 2-42 nm on 485 days, and identified 236 NPF events on 213 days. The majority of these events (37%) occurred during the daylight hours with just 10% at night. However, the NAIS also showed particle growth with no NPF on many nights (28%). Using a scanning mobility particle sizer (SMPS), we showed that particle growth continued at larger sizes and occurred on 70% of nights, typically under high relative humidities. Most particles in the air, especially near coastal locations, contain hygroscopic salts such as sodium chloride that may exhibit deliquescence when the relative humidity exceeds about 75%. The growth rates of particles at night often exceeded the rates observed during NPF events. Although most of these night time growth events were preceded by daytime NPF events, the latter was not a prerequisite for growth. We conclude that particle growth in the atmosphere can be easily misidentified as NPF, especially when they are monitored by an instrument that cannot detect them at the very small sizes.

42
43
44

Keywords: New particle formation, particle growth, atmospheric aerosols, secondary particles.

45

46 **1 Introduction**

47 The formation of secondary particles in the atmosphere through homogeneous nucleation is
48 known as new particle formation (NPF). This is one of the major sources of particles in the
49 atmosphere. The condensable species that contribute are mainly sulfuric acid and semivolatile
50 organic compounds and the process is thought to occur by binary water-sulfuric-acid or
51 ternary water-sulfuric-acid-ammonia nucleation. Particles, thus formed, form stable clusters
52 that continue to grow to larger sized particles by vapour condensation or by coagulation with
53 other particles (Kulmala et al., 2013).

54

55 The particle formation rate and the particle growth rate are the two most important
56 parameters used to characterize an NPF event. The particle formation rate is the rate of
57 formation of smallest measurable size of the particles, generally in the size range 2-3 nm.
58 This is different to the actual nucleation rate (the rate at which the stable clusters form). The
59 particle growth rate varies with particle size (Manninen et al., 2010;Gagné et al.,
60 2011;Backman et al., 2012) and, hence, the reported values depend on the detectable size
61 ranges of the instruments used. Until recently, studies have been limited to measure the
62 particles above 3 nm. However, it is only during the past decade that the advancement of
63 instruments has developed to such a level that particles of 2 nm or even smaller can be
64 measured (Kulmala et al., 2012).

65

66 NPF has been observed under a range of environmental conditions, on every continent in the
67 world (Kulmala et al., 2004;Backman et al., 2012;Gagné et al., 2011;Manninen et al.,
68 2009;Manninen et al., 2010;Rose et al., 2015;Pushpawela et al., 2018;Jayaratne et al., 2017).
69 The occurrence rate of NPF is mainly dependent on the nature and concentration of gaseous

70 precursors, which are controlled by a number of factors including the type and intensity of the
71 sources, concentration of pre-existing aerosols, origin of air masses, photo-chemical
72 processes and meteorological parameters such as intensity of solar radiation, temperature,
73 relative humidity, wind direction and wind speed (Birmili and Wiedensohler, 2000;Kulmala
74 et al., 2004;Kulmala et al., 2013). Pre-existing aerosols act as sinks to condensable gases that
75 are present in the atmosphere. This leads to a reduction in their vapour pressure and inhibits
76 homogeneous nucleation.

77

78 Oxides of nitrogen and volatile organic compounds are readily produced in urban
79 environments from sources such as motor vehicles and industrial facilities (Seinfeld and
80 Pandis, 2006;Harrison, 2007). These gases react with ozone in the presence of sunlight to
81 produce OH radicals that can oxidise gaseous precursors such as sulphur dioxide and nitric
82 oxide, converting them into the condensable species sulfuric acid and nitric acid,
83 respectively. These photochemical reactions are more likely to occur during the day time on
84 sunny days with high intensity of solar radiation, which is when we would expect to observe
85 more NPF events.

86

87 Numerous studies in many different environments have conclusively shown that the large
88 majority of NPF occur during the day time (Seinfeld and Pandis, 2006;Suni et al., 2008;Man
89 et al., 2015;Pushpawela et al., 2018). Very few studies have reported the occurrence of NPF
90 during the night time and these have mostly been in forest environments and coastal sites.
91 Table 1 gives a summary of studies in chronological order, that have reported observations
92 and frequencies of occurrence of night time NPF events, together with the respective
93 frequencies of occurrence of day time NPF events and the instrumentation that was used. We

94 see that, at a given location, NPF events were generally more likely to occur during the day
95 time than during the night. The sole exception is the short study of 16 days by Kammer et al.
96 (2017). Night time events were reported on between 4% and 37% of the days observed. They
97 were more likely to be observed at forest locations (16% to 37%), while the two studies
98 conducted at coastal locations showed significantly lower values of 4% and 11%. In a
99 previous study carried out in and around Brisbane with an SMPS, Salimi et al. (2017)
100 reported NPF events on around one in every four nights. They also reported NPF on every
101 second day which is significantly higher than any of the values found in Brisbane (Guo et al.,
102 2008;Cheung et al., 2011;Crilley et al., 2014;Jayaratne et al., 2016;Pushpawela et al., 2018).

103

104 In the present study, we collected data of charged and uncharged particle concentrations in
105 the urban environment of Brisbane using a neutral cluster and air ion spectrometer (NAIS) on
106 close to five hundred days. The NAIS can provide more accurate information on NPF than
107 the SMPS, because of its ability to measure particles down to 2 nm in size, which is very
108 close to the size at which the initial steps of nucleation and formation of particles occur
109 (Manninen et al., 2011;Manninen et al., 2016). The results were compared with that obtained
110 simultaneously with an SMPS with a minimum detectable size of 9 nm. The SMPS data were
111 also used to determine the growth rates of particles. The observations by the NAIS and SMPS
112 were used to differentiate between (a) local NPF events followed by particle growth and (b)
113 growth events in the absence of NPF events – two phenomena that are not always concurrent
114 and often misidentified when only one instrument is used.

115

116 **2 Methods**

117

118 **2.1 Monitoring Site**

119 The instruments were housed in a sixth-floor laboratory in a building at the Gardens Point
120 campus of the Queensland University of Technology in Brisbane, Australia. The site is
121 situated at the edge of the Brisbane Central Business District bordered by the City Botanical
122 Gardens and the Brisbane River, approximately 100 m away from a busy motorway carrying
123 about 120,000 vehicles per day and is representative of a typical urban environment in
124 Australia. The measurements were carried out during the three calendar years 2012, 2015 and
125 2017, yielding 485 complete days of data.

126 The pollutants at this site were mainly from motor vehicle exhaust emissions. Depending on
127 the wind direction, emissions may also be received from the Port of Brisbane and two oil
128 refineries in its vicinity as well as from Brisbane Airport, all located about 20 km to the
129 north-east of the monitoring site.

130 Meteorological data such as temperature, relative humidity, solar radiation, rainfall, wind
131 direction and wind speed as well as air quality data such as sulphur dioxide (SO₂), ozone
132 (O₃), PM₁₀, PM_{2.5} and atmospheric visibility were obtained from the Department of
133 Environmental and Heritage Protection, Queensland, at their in-situ site at the Queensland
134 University of Technology and two other sites within a distance of 1.5 km from the
135 University.

136

137 **2.2 Description of the instruments**

138 The NAIS, manufactured by Airel Ltd, Estonia (Manninen et al., 2016), detects the mobility
139 distribution of charged clusters and particles of both polarities in the electrical mobility range
140 from 3.2 to 0.0013 cm² V⁻¹s⁻¹. It also measures the size distribution of total particles in the
141 size range from 2.0 - 42 nm. The instrument has a high-resolution time down to 1 s and
142 consists of two cylindrical electrical mobility analysers, one for each polarity. It operates in

143 four modes: ion mode; particle mode; alternate charging mode and offset mode. In the ion
144 mode, the NAIS measures naturally charged particles without any modification. In the
145 particle mode, it uses a corona needle to charge the particles. This leads to an inherent
146 problem where the very small particles cannot be distinguished from the corona ions
147 (Manninen et al., 2016). For this reason, we have restricted the lower detection limit in the
148 particle mode to 2 nm . The alternate charging mode is similar to the particle mode, but it
149 electrically neutralizes the sampled particles and improves the performance of the instrument.
150 In the offset mode, the NAIS measures zero signals, noise levels and parasitic currents. The
151 measurement process of the instrument is fully automated. The measurement cycle of the
152 NAIS varies from 2-5 minutes. A more detailed discussion of its design and principles is
153 given in (Manninen et al., 2011) and (Mirme and Mirme, 2013). In this study, we set the
154 measurement cycle to 2 min ion mode, 2 min particle mode, and 1 min offset mode.

155 An SMPS, consisting of a TSI model 3071 differential mobility analyser and a TSI model
156 3782 condensation particle counter, was used to measure the particle size distribution in the
157 range from 9 - 415 nm.

158

159 **2.3 Data Analysis**

160 **2.3.1 Classification of New Particle Formation (NPF) events:**

161 We identified NPF events using the rate of change of total particle concentration, dN/dt ,
162 where N is the number of particles in the size range 2.0 -10.0 nm and using the classification
163 described by (Zhang et al., 2004). Events with $N > 10,000 \text{ cm}^{-3}$ for at least 1 hour and dN/dt
164 $>10,000 \text{ cm}^{-3}\text{h}^{-1}$ were defined as “strong” NPF events. Events with $5000 < N < 10,000 \text{ cm}^{-3}$
165 for at least 1 hour and $5000 < dN/dt < 10,000 \text{ cm}^{-3}\text{h}^{-1}$ were classified as “weak” NPF events.
166 All of these events started in the nucleation mode size range and prevailed over a time span

167 of more than one hour, generally exhibiting a “banana” shape in the time-series contour plot
168 of particle number concentration (PNC), indicating particle formation and subsequent growth.
169 A 24-hour day that included at least one NPF event was labelled as an ‘NPF Day’. A day on
170 which there were no NPF events was labelled as a ‘Non-event Day’.

171 Every NPF event was characterised by a sharp increase of the PNC in the intermediate size
172 range from 2.0-7.0 nm. This observation has been used to determine the starting time of an
173 NPF event (Leino et al., 2016). Similarly, in the present study, the starting time of a strong
174 NPF event was determined by noting the time of first occurrence of $dN/dt > 10,000 \text{ cm}^{-3}\text{h}^{-1}$.
175 The starting time of a weak NPF event was determined by noting the time of first occurrence
176 of $dN/dt > 5000 \text{ cm}^{-3}\text{h}^{-1}$. N is the number of particles in the size range 2.0-10 nm.

177 NPF events that started between sunrise and sunset were categorized as “day time” NPF. NPF
178 events that started between sunset and sunrise were categorized as “night time” NPF.

179

180 **2.3.2 Classification of Growth events:**

181 The data from the NAIS showed that growth events were not always preceded by an NPF
182 event. Growth events that did not follow an NPF appeared as a “floating-banana” shape in the
183 PNC contour plots. These events were identified using the rate of change in the diameter (d_p)
184 of particle, dd_p/dt . Events with $dd_p/dt > 1 \text{ nm h}^{-1}$ were classified as “growth” events. In the
185 NAIS data, these events showed an enhancement of PNC in the size range above 7 nm.
186 Further, in these events, unlike in NPF events, the sharp increase in PNC in the size range
187 between 2-7 nm was absent. In this way, growth events could be clearly distinguished from
188 NPF events. In fact, unless they were preceded by an NPF event, most growth events showed
189 very few particles in the size range below 10 nm. We also observed “vertical band” shapes
190 which were due to the sudden appearance of high concentrations of particles in all sizes.

191 These were neither NPF nor growth events and characterised the influx of already formed
192 particles from further locations to the monitoring site, and were ignored in the analysis.

193

194 **2.3.3 Calculation of particles growth rate**

195 The growth rate (GR) of particles is defined as

$$196 \quad GR = \frac{dd_p}{dt} = \frac{d_{p2} - d_{p1}}{t_2 - t_1} \quad (1)$$

197 where dp_2 and dp_1 are the diameters of particles at times t_1 and t_2 . This was calculated by the
198 maximum concentration method described in (Kulmala et al., 2012). The unit of the GR is
199 nanometres per hour. During an NPF or a growth event, the number concentration of small
200 particles increases, showing a peak in the particle size distribution. When the particles grow
201 in size, this peak shifts towards larger sizes. In order to derive the maximum particle
202 concentration, we plotted the time series of the PNC in different size ranges. We estimated
203 the GR from the slope of the best-fitted line on the graph of mid-point diameter of particles
204 versus the time of maximum concentration (Dos Santos et al., 2015;Pierce et al., 2014).

205 **2.3.4 Statistically significant differences**

206 Statistical significances of the difference between two parameters were calculated using the
207 Student's t test.

208

209

210 **3. Results and Discussion**

211 **3.1 Observation of NPF during study period**

212

213 The study yielded complete 24h data on a total of 485 days. The instrument was unavailable
214 on some days, as it was required for other projects or was being serviced or cleaned. In
215 addition, a few days were ‘lost’ due to missing data owing to power failures or instrument
216 malfunction. A summary of the observational periods, together with the corresponding
217 number of days on which 24h data were available and NPF events were observed, is shown in
218 Table 2. Columns 3 to 8 represent the number of day time, night time and total NPF classified
219 into strong and weak events according to the method described in section 2.3.1. The last three
220 columns give a summary of all NPF events.

221

222 Altogether, 236 NPF events (strong and weak) were observed on 213 of the 485 days on
223 which we were able to obtain data. Out of this, strong NPF events were observed on 177
224 days, giving an occurrence rate of 37%. This is only slightly less than the rate of 41% found
225 by Pushpawela et al. (2018) using the NAIS in Brisbane over the single calendar year 2012.
226 In the two other studies using the NAIS in Brisbane, Crilley et al. (2014) and Jayaratne et al.
227 (2016) reported higher values of 56% and 45% respectively. However, both these previous
228 studies used a slightly different criteria to identify NPF events, that is they excluded the
229 requirement of $N > 10,000 \text{ cm}^{-3}$ for a period of at least 1 hour. The Crilley et al. (2014) study
230 was also conducted over a much shorter period of 36 days only. Table 2 also shows that,
231 although “strong” day time NPF events were observed on 159 days (33%), “strong” night
232 time NPFs were relatively scarce, occurring on just 18 days (4%). Further, “weak” NPF
233 events were observed on 59 days (12%) and these were almost equally distributed between

234 night and day times. Taking into account all strong and weak NPF, day time NPF occurred
235 on 37% of the days while night time NPF occurred on only 10%. In Table 2, it should be
236 noted that a given day may sometimes have both a day time and a night time event. There
237 were 23 such days. In addition, there were 8 days that had two daytime events and no
238 instances of two events during the same night. There have been three previous studies that
239 have used an SMPS to study NPF in Brisbane. Together with the occurrence rates in
240 parenthesis, these were Guo et al. (2008) (35%), Cheung et al. (2011) (26%) and Salimi et al.
241 (2017) (77%).

242

243 **3.2 Diurnal variation**

244

245 Figures 1 (a) and (b) shows the summary of starting times of all NPF events during the day
246 time and night time, respectively, estimated by using the method described in Section 2.3.1.
247 The histograms show the number of events observed in each 30 min period after sunrise and
248 sunset, respectively. The times indicated on the x-axis refer to the end of each 30 min period.
249 In Figure 1(a), the three bars at the extreme left correspond to times before sunrise. We have
250 classified these as night time events. Both of these figures show that most NPF events (71%)
251 began during the morning, with a high likelihood of occurrence between 2 and 4 hours after
252 sunrise, corresponding to approximately between 8.00 am and 10.00 am. In particular, 90 out
253 of 236 events occurred during this 2-hour period. This is likely to be a result of several
254 factors such as the higher concentration of precursor gases from motor vehicles during the
255 morning rush hour and the onset of solar radiation. However, no NPF were observed during
256 the evening rush hour period around 4-6 pm. During this time, the air temperatures are still
257 relatively high and, although the gaseous precursors are being produced, the vapour pressures
258 may not be sufficiently high to produce secondary particles. The starting times of night time

259 NPF events also showed a distinct trend with a peak likelihood between 3 and 4 hours after
260 sunset, corresponding to approximately 8 and 9 pm. By this time of the day, the temperatures
261 have generally fallen sufficiently for vapour pressures to increase. No night time events were
262 observed at all during the second half of the night, between 11 pm and 4 am. Although the
263 temperatures are low during this time, there is minimum production of precursor gases.

264

265 **3.3 Effect of atmospheric parameters**

266

267 A summary of the mean and range of various meteorological and air quality parameters
268 during NPF and non-event days is shown in Table 3. The mean solar radiation intensity on
269 NPF days were significantly higher compared to the other days with mean values of 505 W
270 m^{-2} and 397 W m^{-2} , respectively. Conversely, the mean relative humidity on NPF days was
271 significantly less than on other days with values of 54% and 66%, respectively. The mean
272 relative humidity on NPF days were 59% and 52% during winter and summer months.
273 Therefore, NPF events were more likely to occur on days with low relative humidity and
274 high solar radiation. Similar observations have been reported from several other urban cities
275 such as Melpitz, Germany (Birmili and Wiedensohler, 2000), San Pietro Capofiume, Italy
276 (Hamed et al., 2007) and Pune, India (Kanawade et al., 2014).

277

278 The wind direction on NPF days was mainly from the south to southwest directions, with a
279 mean wind speed of around 1.4 m s^{-1} . The mean air temperature was 17⁰C and 24⁰C on NPF
280 days during winter and summer months. We did not detect any clear differences in wind
281 direction, wind speed and air temperature between NPF days and the other days. In general,
282 most of the NPF events occurred on days when there was no rainfall observed. However, a
283 clear dependence was found between NPF occurrence and atmospheric visibility. The

284 visibility was expressed through the particle back scatter coefficient (BSP) in units of Mm^{-1} .
285 These two parameters are inversely proportional to each other. The BSP observed at 8 am on
286 NPF days was significantly lower on NPF days than on other days, with mean values of 18
287 Mm^{-1} and 31 Mm^{-1} , respectively. A good discussion about the relationship between the
288 occurrence of NPF in Brisbane and the values of BSP may be found in Jayaratne et al.
289 (2015). This study also found that, no NPF events occurred on days when the mean $\text{PM}_{2.5}$
290 exceeded $20 \mu\text{g m}^{-3}$ in Brisbane.

291
292 The presence of high concentration of O_3 under high solar radiation increases the production
293 of OH radicals, and the presence of high concentration of both SO_2 and OH radicals give rise
294 to increased production of H_2SO_4 leading to NPF (Seinfeld and Pandis, 2006; Lee et al.,
295 2008). Therefore, we would expect SO_2 and O_3 concentration levels to be higher on NPF
296 days than on non-event days. However, we observed only a marginal increase of SO_2 and O_3
297 concentrations on NPF days (Table 3).

298

299 **3.4 Day time and night time NPF events**

300

301 The two upper panels in Figure 2 show the NAIS spectragrams obtained between 8:00 am
302 and 4:00 pm on 19 August 2017 and 31 July 2015, respectively. On 19 August, a strong NPF
303 event began in the morning at around 9:00 am and lasted for 4-5 hours. Here, the total PNC
304 increased from about $30,000 \text{ cm}^{-3}$ at 9:00 am to just over $90,000 \text{ cm}^{-3}$ at 11:00 am, giving a
305 particle formation rate of $30,000 \text{ cm}^{-3} \text{ h}^{-1}$. Thereafter, particles continued to grow in size for
306 several hours. The PNC decreased gradually in the afternoon. The particles showed a
307 relatively high growth rate of about 7 nm h^{-1} in the size range 2-42 nm.

308 The two lower panels in Figure 2 show NAIS spectragrams obtained during the night,
309 between 6:00 pm and 2:00 am on 20 August 2015 and 5 September 2015, respectively. On 20
310 August, a strong NPF event began in the night at around 9:30 pm and lasted for 2-3 hours.
311 The particles also showed a relatively high growth rate of about 11 nm h^{-1} in the size range 2-
312 42 nm.

313
314 We did not observe a significant difference in growth rates of particles between daytime and
315 night time NPF events. The growth rate of particles in the size range 2-42 nm during all NPF
316 events, calculated from equation (1), varied between 4 nm h^{-1} and 22 nm h^{-1} with a mean and
317 standard deviation of $(12.1 \pm 6.5) \text{ nm h}^{-1}$. These growth rates were comparable to the values
318 reported at two other urban locations; Atlanta, USA ($3\text{-}20 \text{ nm h}^{-1}$)(Stolzenburg et al., 2005)
319 and Budapest, Hungary ($2\text{-}13 \text{ nm h}^{-1}$) (Salma et al., 2011). However, the mean values of
320 growth rates obtained by previous studies in Brisbane were significantly lower than the value
321 reported by this study. For example, Cheung et al. (2011) and Salimi et al. (2017) reported
322 growth rates of 4.6 nm h^{-1} and 2.4 nm h^{-1} respectively. Both these studies were carried out
323 using an SMPS with a lower detection size of about 10 nm.

324
325 Typically, the particle growth rates were high during the first few hours and then decreased to
326 a few nanometres per hour within 3-4 hours after nucleation. Several studies have reported
327 that the growth rate of particles in the size range 7-20 nm was greater than that in the smaller
328 size range 3-7 nm (Manninen et al., 2010;Gagné et al., 2011;Yli-Juuti et al., 2009;Backman et
329 al., 2012). Manninen et al. (2010) studied NPF events at 12 European sites and found that 9
330 out of the 12 sites showed this trend while at 3 sites the growth rate was greater in the smaller
331 size range. They suggested that this size dependence was due to different condensing vapours
332 participating in the growth of different sized particles depending on their saturation vapour
333 pressures. For example, it is well known that sulfuric acid plays a dominant role in nucleation

334 and the initial growth of particles during NPF while organics dominate the growth at larger
335 sizes of 10-30 nm (Yli-Juuti et al., 2011;Manninen et al., 2009;Smith et al., 2008). Further
336 evidence comes from the observation that the growth rate of the particles in the larger size
337 range of 7-20 nm is enhanced during the summer when the concentration of biogenic volatile
338 organic compounds in the atmosphere is greater (Yli-Juuti et al., 2011). Our observations of
339 particle growth rates in the different size ranges agree with previous studies that have
340 suggested that the dominant condensable vapour in Brisbane is probably sulfuric acid, with
341 organics playing a secondary role (Crilley et al., 2014).

342

343 **3.5 Observations of growth events during the study period**

344

345 NPF events are almost always followed by particle growth. However, with the NAIS, we
346 observed several growth events that were not preceded by an NPF event. These events were
347 observed more often at night than during the day. A summary of these events observed by the
348 NAIS, is shown in Table 4. Columns 3 to 5 represent the number of day time, night time and
349 total growth events classified according to the method described in section 2.3.2. Figure 3
350 shows examples of NAIS spectragrams of such growth events that occurred during the day
351 time (a) and night time (b). Particle growth is again demonstrated by the typical banana shape
352 of the colour contours, with the difference that the lower end of the ‘banana’ does not reach
353 as far as the smallest particle sizes, indicating that there is no NPF. This shape is sometimes
354 referred to as a “floating banana”, to differentiate it from the complete “banana” shape of an
355 NPF event. In most of the events, particle growth is observed to continue for several hours.
356 The observed rates of growth varied between 1 nm h^{-1} and 45 nm h^{-1} with a mean and
357 standard deviation of $(16.8 \pm 11.9) \text{ nm h}^{-1}$ in the size range 8-42 nm. During the 485 days of
358 observation, excluding NPF events, day time growth events were observed on just 54 days

359 (11%), whereas night time growth events were observed on 135 days (28%). The overall
360 occurrence rate of growth events obtained by the NAIS was 37%. However, it should be
361 noted that particles continued to grow at sizes larger than the upper size detection rate of the
362 NAIS, which was 42 nm. Thus, the SMPS was likely to detect many more growth events than
363 the NAIS.

364

365 **3.6 Observations of particle growth by SMPS**

366

367 Next, we look at the behaviour of total PNC and the median particle diameter of NPF and
368 growth events using the data obtained by the SMPS. Figure 4 shows a period of 6 days,
369 during which there were 3 consecutive daytime NPF events that were followed by two non-
370 event days and a day with a daytime NPF event. The NPF events are shown by red arrows. In
371 each of these four cases, prior to the inception of the daytime NPF, the total PNC was low -
372 about 2500 cm^{-3} . During the NPF event, the total PNC increased from about 5000 cm^{-3} in the
373 morning to over $15,000 \text{ cm}^{-3}$ near mid-day. Thereafter, the particles started to grow in size up
374 to 20-30 nm. During and after the late afternoon, although the total PNC began to decrease,
375 the particles continued to grow in size up to 40-65 nm. All 4 NPF events continued through
376 this “second phase of particle growth” until the early hours of the next day. The growth rate
377 varied between $2\text{-}7 \text{ nm h}^{-1}$.

378 Figure 5 shows another example. During this 7 day period, two growth events in the late
379 afternoon were preceded by NPF events. The remaining two growth events did not follow
380 any NPF event. The particles grew up to 40-50 nm. During the measurement period, particle
381 growth events were observed on 65-70% of the nights.

382

383 These observations of continued growth of particles following NPF events is a common
384 phenomenon and has been reported by several other researchers. For example, Man et al.
385 (2015) observed 12 out of 17 NPF events with particle growth from 10 nm to 40 nm during
386 the day time at a suburban coastal site in Hong Kong. In addition, they observed 3 events
387 with second phase of particle growth to 61-97 nm at night time. These three events were
388 preceded by a daytime NPF event. Russell et al. (2007) observed nanoparticle growth on 19
389 out of 48 days (40%) during the day time and on 5 out of 48 days (10%) during the night time
390 in Appledore Island, Maine, USA. Subsequently, particle growth continued over several
391 hours with rates varying from 3 to 13 nm h⁻¹.

392

393 NPF generally occur at high solar radiation, high temperature and low relative humidity.
394 However, growth events were more likely to occur during time periods with low
395 temperature and high relative humidity. We investigated this further by plotting the median
396 particle size and relative humidity as a function of time during growth events (Figure 6). In
397 general, progression into the night time, after 6:00 pm, was accompanied by a decrease in air
398 temperature, resulting in an increase in relative humidity in the atmosphere.

399

400 During the event that occurred on July 16, 2012 the median particle size increased from about
401 30 to 65 nm as the relative humidity increased from 65% to 80% (Figure 6a). Similarly,
402 during the event that occurred on July 20, 2012 the median particle size increased from about
403 30 to 75 nm as the relative humidity increased from 75% to 90% (Figure 6b).

404

405 It is well-known that relative humidity may favour particle growth in the atmosphere owing
406 to several reasons. For example, atmospheric aerosol particles increase in size with relative
407 humidity due to the uptake of water (Winkler, 1988). In addition, when the relative humidity

408 increases, heterogeneous reactions can take place in the liquid phase of a growing particle
409 while, if there is an accompanying drop in temperature, it would enhance the transport of
410 semivolatile compounds from the gas phase on to the surface of the particles. Water uptake is
411 caused by the deliquescence of soluble salts which form a solution when the solid compound
412 is exposed to water vapour at sufficiently high vapour pressure. Several organic materials are
413 also known to absorb water at high humidity which is more generally known as
414 hygroscopicity. Sodium chloride (NaCl) has a deliquescence point of 76% relative humidity.
415 At this point, a NaCl-bearing particle will deliquesce and become a solution of droplet with a
416 well-defined spherical shape. The particle diameter does not change considerably as the
417 relative humidity is increased from 0 to 74%, beyond which it can increase considerably.
418 Close to the coast, sea-salt aerosols constitute a large proportion of the atmospheric
419 particulate mass and NaCl is a major component. Many of the inorganic substances that
420 readily absorb water, such as sea salt, ammonium salts and nitrates, are present in the
421 Brisbane environment (Harrison, 2007). Therefore, it is not surprising that, in the present
422 study, we observed that particle growth occurred on 7 out of 10 nights with high relative
423 humidity.

424

425 **3.7 Probability of growth events being misidentified as NPF events**

426 In Figure 3 (a), the horizontal white line indicates the typical lower size detection threshold of
427 the SMPS that has been used in many locations before; we chose 7 nm as a typical value in
428 this case. The SMPS does not 'see' any particles below this line. It is clear that there is an
429 enhancement of PNC in the size range 7-20 nm around 11:30 am on this day. The absence of
430 intermediate size particles (between 2-7 nm) suggests that the 7-20 nm particles originated
431 on-site by primary emission or were advected to the site from a distant location. The NAIS
432 clearly shows that this was not an NPF event. However, in the absence of information below

433 a particle size of 7 nm, the SMPS data may be easily misinterpreted as an NPF event. The
434 typical ‘floating banana’ shape of the spectrogram contours show that the particles continue
435 to grow between 11:30 a.m. and about 1:00 p.m. and this can be observed by an SMPS. As
436 we have demonstrated, growth events are not always formation events. There are two
437 enhancement events near 1.00 pm and 3.30 pm. Once again, the NAIS shows that neither of
438 these are NPF events, although based on the SMPS they may be mistakenly identified as
439 such. Figure 3 (b) shows another event that can be easily misidentified as an NPF event based
440 on SMPS data alone.

441

442 Salimi et al. (2017), using an SMPS with a lower size limit of 9 nm at 25 sites across
443 Brisbane, reported 219 NPF events out of 285 days of measurements. This occurrence rate of
444 77% (67% of day time and 33% of night time) is significantly higher than any of the values
445 found previously in Brisbane and at any other location in the world. With the NAIS, it was
446 possible to show that most of these events were growth events and not NPF events. It was not
447 possible to differentiate these two types of events with the SMPS alone as it provides no
448 knowledge of the PNC below 9 nm. With the NAIS, we did not observe nocturnal NPF
449 events on more than 47 of 500 days.

450 In many NPF events, particle growth ceases after they have grown to a certain size. In the
451 growth event in Figure 3, the maximum size is about 25 nm. In such cases, the greater part of
452 the ‘banana’ profile is below 7 nm and, thus, invisible to the SMPS. This could result in the
453 missing of such NPF events. Considering, all the factors above, it is clear that the NAIS has a
454 distinct advantage over the SMPS in correctly identifying NPF events in the atmosphere.

455

456 **4. Summary and Conclusions**

457 We monitored charged and neutral PNCs in the size range 2-42 nm on nearly 500 days over
458 three calendar years in the urban environment of Brisbane, Australia, using a NAIS. The data
459 were used to differentiate between NPF events and growth events with no NPF. Day time
460 NPF events were observed on 37% of the observational days, with night time events on only
461 10% of the days. NPF events were more likely to occur on days with low relative humidity
462 and high solar radiation. 71% of NPF events occurred during the morning, with the highest
463 probability of occurrence between 2 and 4 hours after sunrise, corresponding to
464 approximately between 8.00 am and 10.00 am. Most of the night time events occurred
465 between 3 and 4 hours after sunset, corresponding to approximately between 8.00 pm and
466 9:00 pm. No night time events were observed between 11.00 pm and 4.00 am. 28% of the
467 particle growth events that occurred at night were not preceded by an NPF event. These
468 events were characterized by high growth rates of up to 45 nm h⁻¹. The SMPS results showed
469 that particle growth continued at larger sizes from ~40 nm to 70 nm and occurred on 70% of
470 nights. Maximum relative humidities were over 80% on most of these nights. These results
471 show that, when particles are monitored by an instrument such as the SMPS that cannot
472 detect them at the very small sizes, particle growth in the atmosphere may be easily
473 misidentified as NPF, leading to an overestimation of the frequency of the latter.

474 **Acknowledgements**

475 We are thankful to the Department of Environmental and Heritage Protection, Queensland,
476 for providing some of the meteorological data used in this study.

477

478 **References:**

- 479 Backman, J., Rizzo, L. V., Hakala, J., Nieminen, T., Manninen, H. E., Morais, F., Aalto, P.
480 P., Siivola, E., Carbone, S., and Hillamo, R.: On the diurnal cycle of urban aerosols, black
481 carbon and the occurrence of new particle formation events in springtime São Paulo, Brazil,
482 *Atmospheric Chemistry and Physics*, 12, 11733-11751, 2012.
- 483 Birmili, W., and Wiedensohler, A.: New particle formation in the continental boundary layer:
484 Meteorological and gas phase parameter influence, *Geophysical Research Letters*, 27, 3325-
485 3328, 2000.
- 486 Cheung, H., Morawska, L., and Ristovski, Z.: Observation of new particle formation in
487 subtropical urban environment, *Atmospheric Chemistry and Physics*, 11, 3823, 2011.
- 488 Crilley, L. R., Jayaratne, E. R., Ayoko, G. A., Miljevic, B., Ristovski, Z., and Morawska, L.:
489 Observations on the formation, growth and chemical composition of aerosols in an urban
490 environment, *Environmental science & technology*, 48, 6588-6596, 2014.
- 491 Dos Santos, V., Herrmann, E., Manninen, H., Hussein, T., Hakala, J., Nieminen, T., Aalto, P.,
492 Merkel, M., Wiedensohler, A., and Kulmala, M.: Variability of air ion concentrations in
493 urban Paris, *Atmospheric Chemistry and Physics*, 15, 13717-13737, 2015.
- 494 Gagné, S., Lehtipalo, K., Manninen, H., Nieminen, T., Schobesberger, S., Franchin, A., Yli-
495 Juuti, T., Boulon, J., Sonntag, A., and Mirme, S.: Intercomparison of air ion spectrometers:
496 an evaluation of results in varying conditions, *Atmospheric Measurement Techniques*, 4, 805-
497 822, 2011.
- 498 Guo, H., Ding, A., Morawska, L., He, C., Ayoko, G., Li, Y. S., and Hung, W. T.: Size
499 distribution and new particle formation in subtropical eastern Australia, *Environ. Chem.*, 5,
500 382-390, 2008.

501 Hamed, A., Joutsensaari, J., Mikkonen, S., Sogacheva, L., Maso, M. D., Kulmala, M.,
502 Cavalli, F., Fuzzi, S., Facchini, M., and Decesari, S.: Nucleation and growth of new particles
503 in Po Valley, Italy, *Atmospheric Chemistry and Physics*, 7, 355-376, 2007.

504 Harrison, R. M.: *Understanding our environment: an introduction to environmental chemistry*
505 *and pollution*, Royal Society of chemistry, 2007.

506 Jayaratne, E. R., Clifford, S., and Morawska, L.: Atmospheric Visibility and PM10 as
507 Indicators of New Particle Formation in an Urban Environment, *Environmental Science &*
508 *Technology*, 49, 12751-12757, 10.1021/acs.est.5b01851, 2015.

509 Jayaratne, E. R., Pushpawela, B., and Morawska, L.: Temporal evolution of charged and
510 neutral nanoparticle concentrations during atmospheric new particle formation events and its
511 implications for ion-induced nucleation, *Frontiers of Environmental Science & Engineering*,
512 10, 13, 2016.

513 Jayaratne, R., Pushpawela, B., He, C., Li, H., Gao, J., Chai, F., and Morawska, L.:
514 Observations of particles at their formation sizes in Beijing, China, *Atmos. Chem. Phys.*, 17,
515 8825-8835, 10.5194/acp-17-8825-2017, 2017.

516 Junninen, H., Hulkkonen, M., Riipinen, I., Nieminen, T., Hirsikko, A., Suni, T., Boy, M.,
517 LEE, S. H., Vana, M., and Tammet, H.: Observations on nocturnal growth of atmospheric
518 clusters, *Tellus B*, 60, 365-371, 2008.

519 Kalivitis, N., Stavroulas, I., Bougiatioti, A., Kouvarakis, G., Gagné, S., Manninen, H.,
520 Kulmala, M., and Mihalopoulos, N.: Night-time enhanced atmospheric ion concentrations in
521 the marine boundary layer, *Atmospheric Chemistry and Physics*, 12, 3627-3638, 2012.

522 Kammer, J., Perraudin, E., Flaud, P.-M., Lamaud, E., Bonnefond, J., and Villenave, E.:
523 Observation of nighttime new particle formation over the French Landes forest, *Science of*
524 *The Total Environment*, 2017.

525 Kanawade, V., Tripathi, S. N., Siingh, D., Gautam, A. S., Srivastava, A. K., Kamra, A. K.,
526 Soni, V. K., and Sethi, V.: Observations of new particle formation at two distinct Indian
527 subcontinental urban locations, *Atmos. Environ.*, 96, 370-379, 2014.

528 Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V.-M., Birmili,
529 W., and McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: a
530 review of observations, *J. Aerosol Sci.*, 35, 143-176, 2004.

531 Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso,
532 M., Aalto, P. P., Junninen, H., and Paasonen, P.: Measurement of the nucleation of
533 atmospheric aerosol particles, *Nature protocols*, 7, 1651-1667, 2012.

534 Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T.,
535 Petäjä, T., Sipilä, M., Schobesberger, S., and Rantala, P.: Direct observations of atmospheric
536 aerosol nucleation, *Science*, 339, 943-946, 2013.

537 Lee, S. H., Young, L. H., Benson, D. R., Suni, T., Kulmala, M., Junninen, H., Campos, T. L.,
538 Rogers, D. C., and Jensen, J.: Observations of nighttime new particle formation in the
539 troposphere, *Journal of Geophysical Research: Atmospheres* (1984–2012), 113, 2008.

540 Leino, K., Nieminen, T., Manninen, H. E., Petäjä, T., Kerminen, V.-M., and Kulmala, M.:
541 Intermediate ions as a strong indicator of new particle formation bursts in a boreal forest,
542 *Boreal Environment Research*, 2016.

543 Man, H., Zhu, Y., Ji, F., Yao, X., Lau, N. T., Li, Y., Lee, B. P., and Chan, C. K.: Comparison
544 of daytime and nighttime new particle growth at the HKUST supersite in Hong Kong,
545 *Environmental science & technology*, 49, 7170-7178, 2015.

546 Manninen, H., Nieminen, T., Riipinen, I., Yli-Juuti, T., Gagné, S., Asmi, E., Aalto, P., Petäjä,
547 T., Kerminen, V.-M., and Kulmala, M.: Charged and total particle formation and growth rates
548 during EUCAARI 2007 campaign in Hyytiälä, *Atmospheric Chemistry and Physics*, 9, 4077-
549 4089, 2009.

550 Manninen, H., Nieminen, T., Asmi, E., Gagné, S., Häkkinen, S., Lehtipalo, K., Aalto, P.,
551 Vana, M., Mirme, A., and Mirme, S.: EUCAARI ion spectrometer measurements at 12
552 European sites—analysis of new particle formation events, *Atmospheric Chemistry and*
553 *Physics*, 10, 7907-7927, 2010.

554 Manninen, H., Franchin, A., Schobesberger, S., Hirsikko, A., Hakala, J., Skromulis, A.,
555 Kangasluoma, J., Ehn, M., Junninen, H., and Mirme, A.: Characterisation of corona-
556 generated ions used in a Neutral cluster and Air Ion Spectrometer (NAIS), *Atmospheric*
557 *Measurement Techniques*, 4, 2767, 2011.

558 Manninen, H. E., Mirme, S., Mirme, A., Petäjä, T., and Kulmala, M.: How to reliably detect
559 molecular clusters and nucleation mode particles with Neutral cluster and Air Ion
560 Spectrometer (NAIS), *Atmos. Meas. Tech. Discuss*, 2016.

561 Mazon, S. B., Kontkanen, J., Manninen, H. E., Nieminen, T., Kerminen, V.-M., and Kulmala,
562 M.: A long-term comparison of nighttime cluster events and daytime ion formation in a
563 boreal forest, *Boreal Environment Research*, 2016.

564 Mirme, S., and Mirme, A.: The mathematical principles and design of the NAIS-a
565 spectrometer for the measurement of cluster ion and nanometer aerosol size distributions,
566 *Atmospheric Measurement Techniques*, 6, 1061, 2013.

567 Pierce, J., Westervelt, D., Atwood, S., Barnes, E., and Leaitch, W.: New-particle formation,
568 growth and climate-relevant particle production in Egbert, Canada: analysis from 1 year of
569 size-distribution observations, *Atmospheric Chemistry and Physics*, 14, 8647-8663, 2014.

570 Pushpawela, B., Jayaratne, R., and Morawska, L.: Temporal distribution and other
571 characteristics of new particle formation events in an urban environment, *Environmental*
572 *Pollution*, 233, 552-560, , 2018.

573 Rose, C., Sellegri, K., Asmi, E., Hervo, M., Freney, E., Colomb, A., Junninen, H., Duplissy,
574 J., Sipilä, M., and Kontkanen, J.: Major contribution of neutral clusters to new particle

575 formation at the interface between the boundary layer and the free troposphere, *Atmospheric*
576 *Chemistry and Physics*, 15, 3413-3428, 2015.

577 Russell, L., Mensah, A., Fischer, E., Sive, B., Varner, R., Keene, W., Stutz, J., and Pszenny,
578 A.: Nanoparticle growth following photochemical α - and β -pinene oxidation at Appledore
579 Island during International Consortium for Research on Transport and
580 Transformation/Chemistry of Halogens at the Isles of Shoals 2004, *Journal of Geophysical*
581 *Research: Atmospheres*, 112, 2007.

582 Salimi, F., Rahman, M., Clifford, S., Ristovski, Z., and Morawska, L.: Nocturnal new particle
583 formation events in urban environments, *Atmospheric Chemistry and Physics*, 17, 521-530,
584 2017.

585 Salma, I., Borsós, T., Weidinger, T., Aalto, P., Hussein, T., Dal Maso, M., and Kulmala, M.:
586 Production, growth and properties of ultrafine atmospheric aerosol particles in an urban
587 environment, *Atmospheric Chemistry and Physics*, 11, 1339, 2011.

588 Smith, J., Dunn, M., VanReken, T., Iida, K., Stolzenburg, M., McMurry, P., and Huey, L.:
589 Chemical composition of atmospheric nanoparticles formed from nucleation in Tecamac,
590 Mexico: Evidence for an important role for organic species in nanoparticle growth,
591 *Geophysical Research Letters*, 35, 2008.

592 Stolzenburg, M. R., McMurry, P. H., Sakurai, H., Smith, J. N., Mauldin, R. L., Eisele, F. L.,
593 and Clement, C. F.: Growth rates of freshly nucleated atmospheric particles in Atlanta,
594 *Journal of Geophysical Research: Atmospheres*, 110, 2005.

595 Suni, T., Kulmala, M., Hirsikko, A., Bergman, T., Laakso, L., Aalto, P., Leuning, R., Cleugh,
596 H., Zegelin, S., and Hughes, D.: Formation and characteristics of ions and charged aerosol
597 particles in a native Australian Eucalypt forest, *Atmospheric Chemistry and Physics*, 8, 129-
598 139, 2008.

599 Svenningsson, B., Arneth, A., Hayward, S., Holst, T., Massling, A., Swietlicki, E., Hirsikko,
600 A., Junninen, H., Riipinen, I., and Vana, M.: Aerosol particle formation events and analysis
601 of high growth rates observed above a subarctic wetland–forest mosaic, *Tellus B*, 60, 353-
602 364, 2008.

603 Winkler, P.: The growth of atmospheric aerosol particles with relative humidity, *Physica*
604 *Scripta*, 37, 223, 1988.

605 Yli-Juuti, T., Riipinen, I., Aalto, P. P., Nieminen, T., Maenhaut, W., Janssens, I. A., Claeys,
606 M., Salma, I., Ocskay, R., and Hoffer, A.: Characteristics of new particle formation events
607 and cluster ions at K-puszta, Hungary, *Boreal Environment Research*, 14, 683-698, 2009.

608 Yli-Juuti, T., Nieminen, T., Hirsikko, A., Aalto, P., Asmi, E., Hörrak, U., Manninen, H.,
609 Patokoski, J., Maso, M. D., and Petäjä, T.: Growth rates of nucleation mode particles in
610 Hyytiälä during 2003–2009: variation with particle size, season, data analysis method and
611 ambient conditions, *Atmospheric Chemistry and Physics*, 11, 12865-12886, 2011.

612 Zhang, Q., Stanier, C. O., Canagaratna, M. R., Jayne, J. T., Worsnop, D. R., Pandis, S. N.,
613 and Jimenez, J. L.: Insights into the chemistry of new particle formation and growth events in
614 Pittsburgh based on aerosol mass spectrometry, *Environmental science & technology*, 38,
615 4797-4809, 2004.

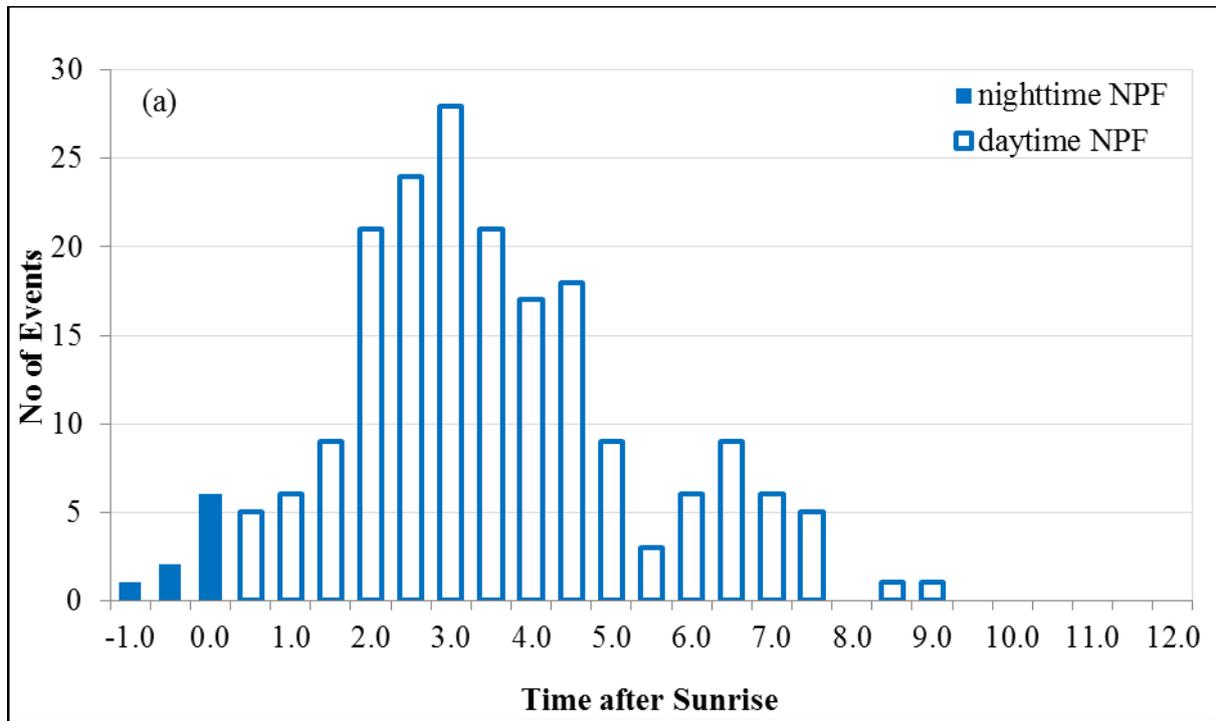
616

617

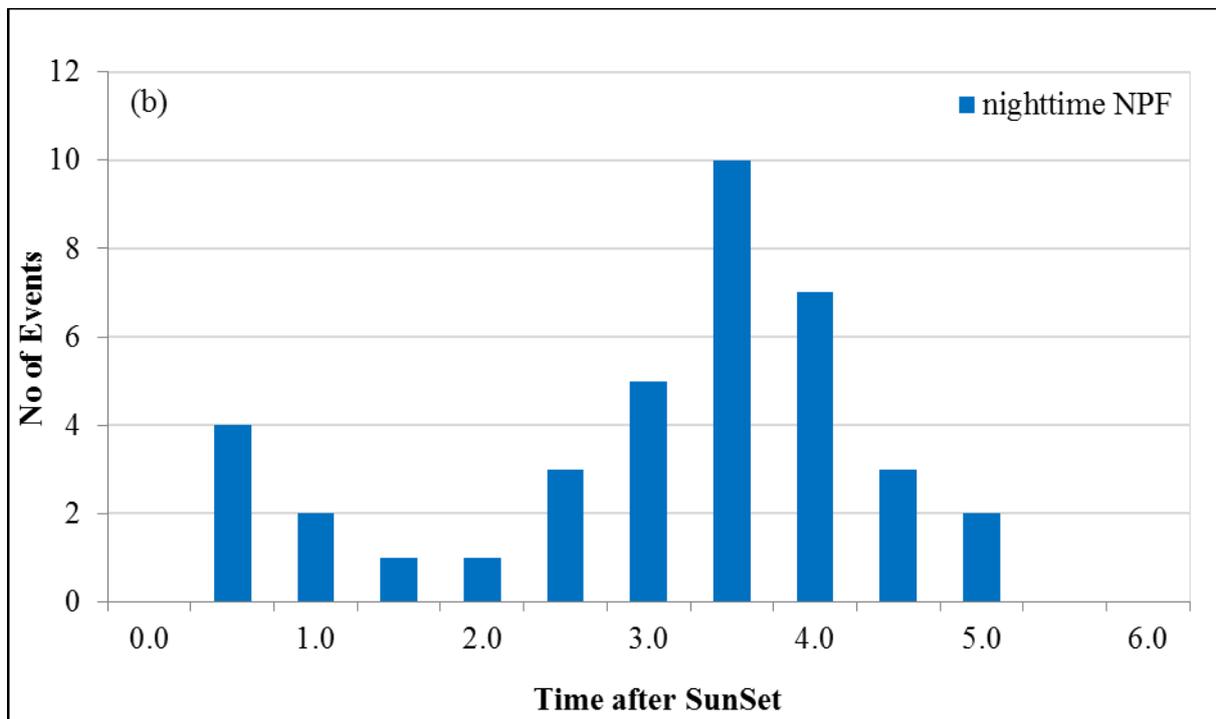
618

Figures

619



620

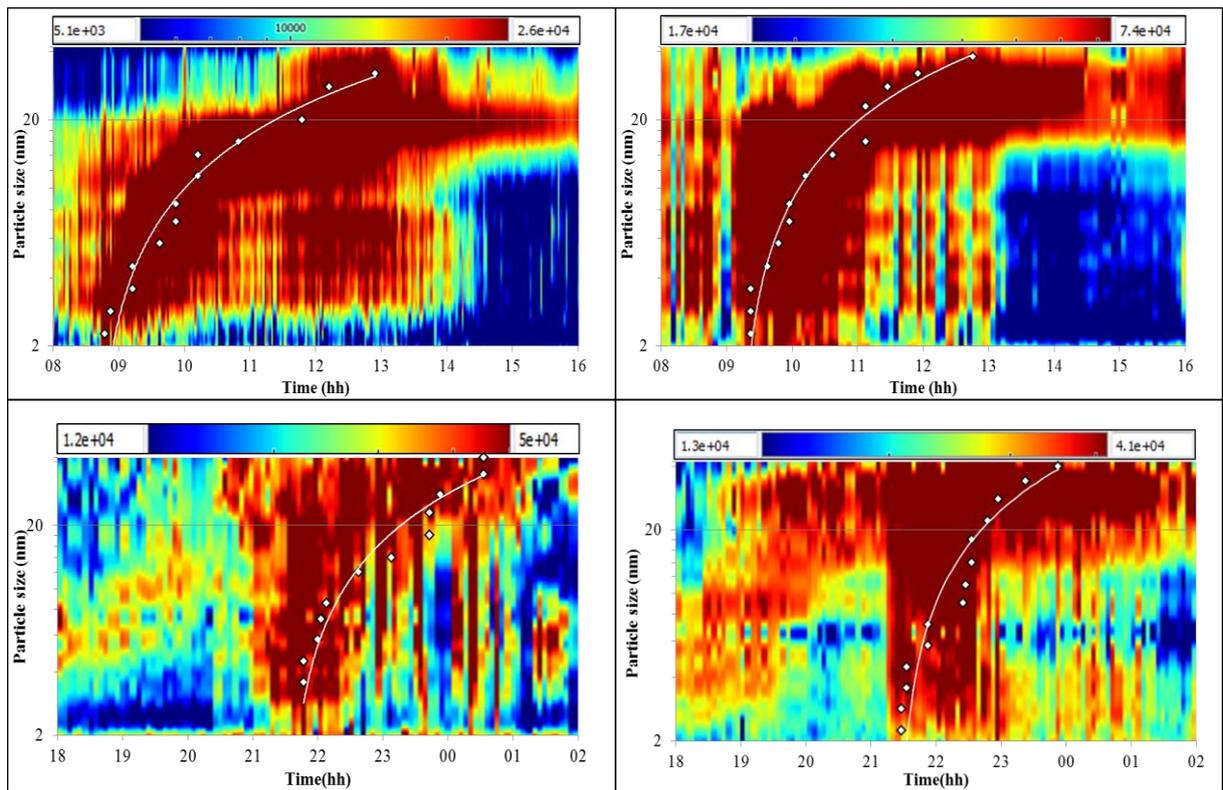


621

622 Figure 1: Distribution of start times of daytime NPF events as a function of time after sunrise

623 (a) nighttime NPF events as a function of time after sunset (b). In figure 1 (a) the three bars

624 on the extreme left correspond to times before sunrise.



626

627 Figure 2: NAIS spectragrams of the daytime NPF events (upper panel) and nighttime NPF

628 (lower panel). The colour contour represents the PNC and the markers represent the times at

629 which the PNC reached its maximum value at each particle size. The unit of PNC is per cubic

630 centimetre.

631

632

633

634

635

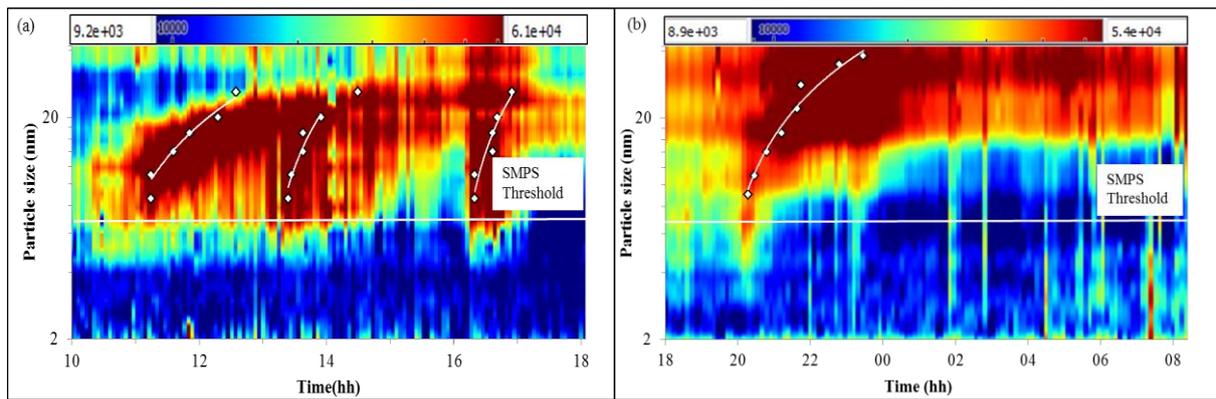
636

637

638

639

640



641

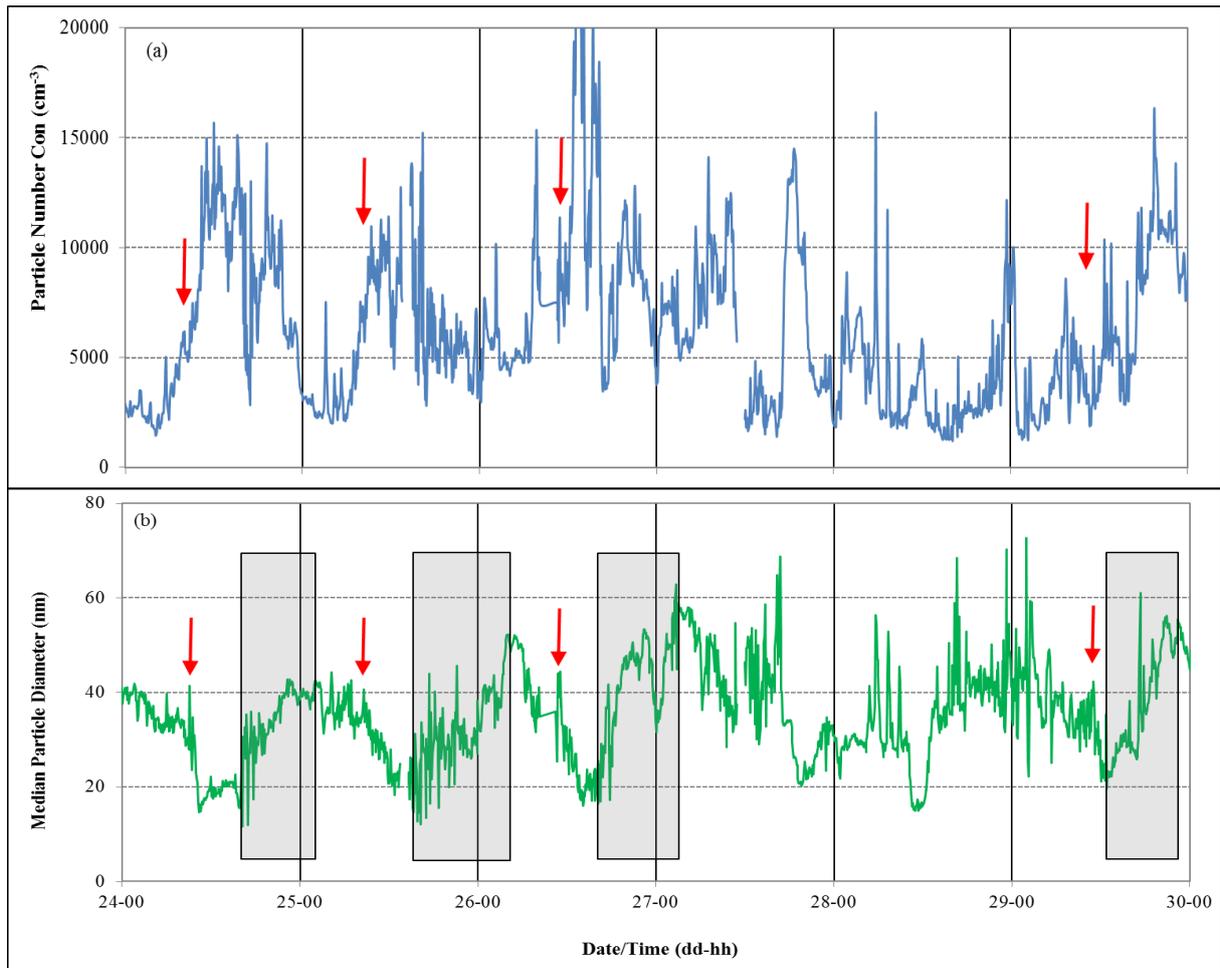
642

643 Figure 3: NAIS spectragrams of the growth events that occurred during (a) day time (b) night
 644 time. Note the “floating banana” shape which indicates that these are clearly not NPF events.

645 The SMPS cannot detect particles at sizes below the horizontal white line. The colour contour
 646 represents the PNC and the markers represent the times at which the PNC reached its
 647 maximum value at each particle size. The unit of PNC is per cubic centimetre.

648

649



650

651

652 Figure 4: (a) the total PNC and (b) median particle diameter from the SMPS during 24 July-

653 30 July, 2012. Red arrows and gray boxes represent the day time NPF events and the growth

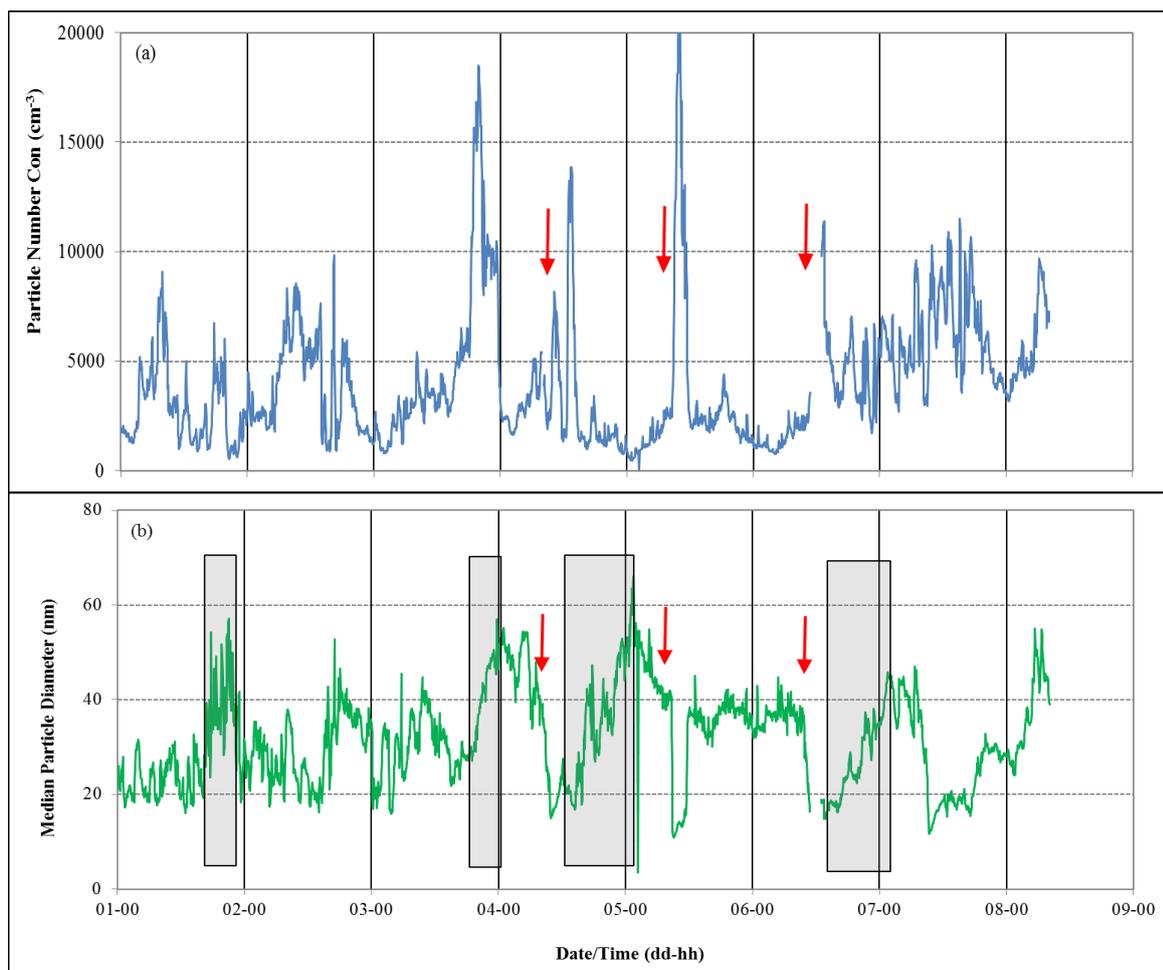
654 events, respectively.

655

656

657

658



659

660

661 Figure 5: (a) the total PNC and (b) median particle diameter from the SMPS during 1 June-7

662 June, 2012. Red arrows and gray boxes represent the daytime NPF events and the growth

663 events, respectively.

664

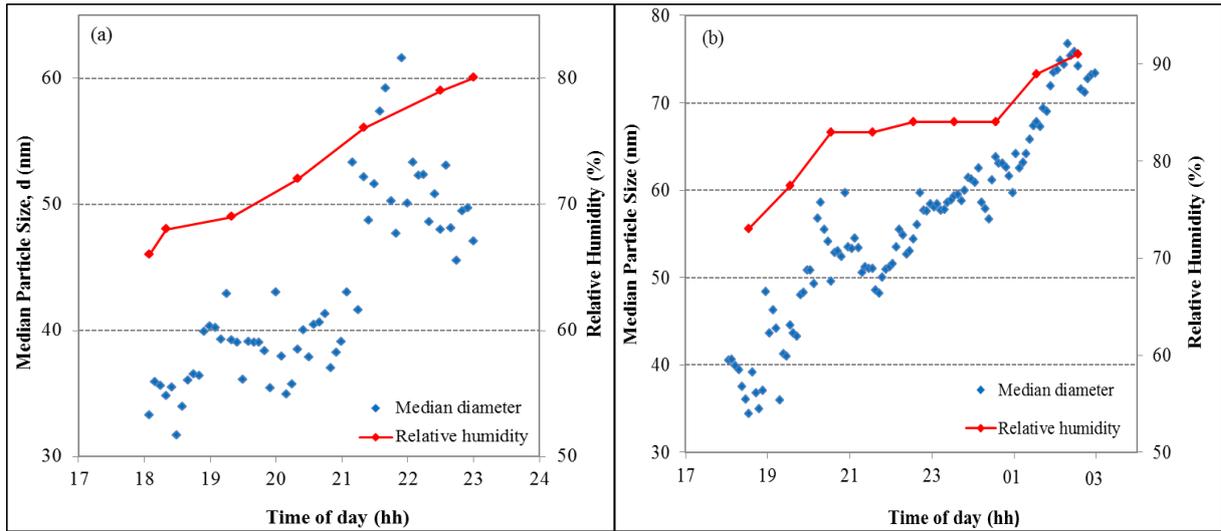
665

666

667

668

669



670

671 Figure 6: Median particle size and relative humidity as a function of time for growth events

672 on July 16 and July 20, 2012, respectively

673

674

675

676

677

678

679

680

681

682

683

684

685

686

687

688

689 **Table 1: Summary of studies reporting night time NPF events**

690 SMPS: Scanning mobility particle sizer, AIS: Air ion spectrometer, BSMA: Balanced

691 scanning mobility analyser, FMPS: Fast mobility particle sizer

Study	Location	Occurrence rate		Instrument (size range)
		Day time	Night time	
Svenningsson et al. (2008)	Abisko, Sweden (characterized by Subartic birch forest)	46/195 days (23%)	31/195 days (16%)	SMPS (10-500 nm) AIS (0.4-40 nm)
Junninen et al. (2008)	Pine Forest, Hyytiala, Finland		344/1279 days (27%)	BSMA (0.4-6.3nm) AIS (0.34-40 nm)
Suni et al. (2008)	Eucalypt forest, Tumbarumba, Australia	184/351 days (52%)	112/351days (32%)	AIS (0.34-40 nm)
Kalivitis et al. (2012)	Finokalia, Lassithiou, Greece (remote coastal site)	53/365 days (15%)	39/365 days (11%)	SMPS (9-900 nm) AIS (0.8-42 nm)
Man et al. (2015)	Suburban coastal site, Hong Kong	12/112 days (11%)	5/112 days (4%)	FMPS (5.6-560 nm)
Mazon et al. (2016)	SMEAR II, boreal forest, Hyytiala, Finland		using neg ions: 1324/4015 days (34%) using pos ions: 1172 /4015 days (30%)	BSMA (0.8-8 nm)
Salimi et al. (2017)	25 sites across Brisbane (characterized by urban environment)	146/285 days (51%)	73/285 days (26%)	SMPS (9-414 nm)
Kammer et al. (2017)	Landes forest, France	2/16 days (12.5%)	6/16 days (37.5%)	SMPS (10-478 nm)

692

693

694 Table 2: Summary of the day time and night time NPF events

695

Year	Total Data Available Days	Strong NPF events			Weak NPF events			Total NPF events		
		Day time	Night time	Total	Day time	Night time	Total	Day time	Night time	Total
2012	253	97	7	104	9	9	18	106	16	122
2015	65	18	4	22	5	7	12	23	11	34
2017	167	44	7	51	16	13	29	60	20	80
Total Events		159	18	177	30	29	59	189	47	236
Total days	485	159	18	177	30	29	59	181	47	213
Occurrence rate (%)		33	4	37	6	6	12	37	10	44

696

697

698

699

700

701

702

703

704

705

706

707

708

709

710

711

712

713

714 Table 3: The mean and the range of meteorology and gas phase parameters on NPF and non-
 715 event days

Parameter	Winter Months		Summer Months		NPF days	non-event days
	NPF days	non-event days	NPF days	non-event days		
Meteorology						
Solar radiation (W m ⁻²)	346 (230-490)	316 (95-476)	600 (202-818)	476 (68-818)	505 (202-818)	397 (68-818)
Temperature (°C)	17 (12-19)	16 (12-25)	24 (18-29)	24 (19-32)	21 (12-29)	20 (12-32)
Relative Humidity (%)	59 (31-73)	70 (27-90)	52 (23-73)	63 (25-86)	54 (23-73)	66 (25-90)
Wind direction (°)	215 S-SW	203 S-SW	197 S-SW	177 S-SW	205 S-SW	200 S-SW
Wind Speed (m s ⁻¹)	1.07 (0.3-3.1)	1.17 (0.3-3.6)	1.60 (0.3-4.7)	2.25 (0.3-5.8)	1.40 (0.3-4.7)	1.72 (0.3-5.8)
Gas Phase						
Visibility (Mm ⁻¹)	15 (6-42)	34 (2-112)	19 (7-41)	29 (6-114)	18 (6-42)	31 (2-114)
Ozone (ppb)	12 (1-29)	10 (2-26)	20 (1-32)	19 (3-35)	17 (1-32)	15 (2-35)
SO ₂ (ppb)	7 (6-10)	6 (1-9)	5 (1-14)	3 (1-9)	6 (1-14)	5 (1-9)

716

717

718

719

720

721

722 Table 4: Summary of the growth events, which did not follow the NPF events, obtained using
 723 the NAIS data.

Year	Total Data Available Days	Growth Events		
		Day time	Night time	Total
2012	253	24	59	83
2015	65	4	21	25
2017	167	26	55	81
Total events		54	135	189
Total days	485	54	135	179
Occurrence rate (%)		11	28	37

724

725

726

727

728