

1 **Author response to referee's comments for acp-**  
2 **2018-159: Production of particulate brown carbon**  
3 **during atmospheric aging of wood-burning**  
4 **emissions by Kumar et al.**

5  
6 Dear editor,

7  
8 We thank both reviewers for their constructive comments, which significantly enhanced the  
9 quality of our manuscript. Below, we provide a point-by-point response (regular typeset) to  
10 the comments (blue font) of both reviewers. The modifications made to the manuscript are in  
11 grey font, indented and *italicized*. Please note that all references to line numbers are to the  
12 submitted manuscript (the ACPD file) and not the revised manuscript.

13  
14 **Anonymous referee 1.**

15 **GENERAL COMMENT**

16 The paper presents an analysis of black carbon (BC) and organic aerosol absorption  
17 properties upon aging. The experiments were done in a laboratory using a smog chamber  
18 where aerosols were exposed to OH radical and UV radiation. Several optical and chemical  
19 properties were measured online and filter samples were collected to be analyzed later offline  
20 by a multiple wavelength absorbance measurement technique and an EC/OC analyzer.  
21 Methanol and water filter extracts were also analyzed and absorbance was measured. The  
22 data is very valuable and the paper is well written and presented. I would recommend its  
23 publication after addressing the comments I present below.

24 My main concerns are related to the following aspects:

25 - The offline techniques used in this study suffer of different artifacts and they are not  
26 sufficiently discussed in the manuscript.

27 We had carefully considered such artifacts, but could have made more reference to them in  
28 the manuscript. We consider the reviewer to be referring to unquantifiable uncertainties.  
29 Based on both reviewers' comments, we have added a section discussing quantifiable and  
30 unquantifiable uncertainties in the method section. This section reads as follows:

31 *Uncertainty analysis. It is important to draw a clear distinction between uncertainties*  
32 *related to measurement precision and accuracy and those related with experimental*  
33 *variability. In this section we discuss the quantifiable and unquantifiable uncertainties*  
34 *related with the different measurements. In the result section, we will present our*  
35 *confidence levels on the average parameters determined based on the experimental*  
36 *variability, which we judge to be the main source of variance in the data.*

37

38 → **Quantifiable uncertainties:**

39 *The estimated uncertainty in the AMS-derived OA mass concentrations is ~25%,*  
40 *which includes both potential biases and precision. This estimate is based on the*  
41 *variation in the AMS calibration factors and estimated uncertainties in the SMPS*  
42 *used for the AMS calibration (Bruns et al., 2015, 2016). Uncertainties related to*  
43 *particle transmission efficiency in the AMS are considered negligible for the particles*  
44 *sampled here (Liu et al., 2007), whose volume size distribution falls within the range*  
45 *transmitted efficiently by the AMS aerodynamic lens (see Fig. S4). The bounce-related*  
46 *collection efficiency (CE) of the AMS was concluded to be unity for wood-burning OA*  
47 *in the literature reviewed by Corbin et al. (2015b; in their Section S1.2). For the*  
48 *present data, the comparison between the SMPS mass (predicted from fitted volume*  
49 *distributions using a density of  $1.5 \text{ g cm}^{-3}$ ) and the total PM predicted as AMS-*  
50 *OA+eBC, suggest a CE value between 0.7 and 1.0 (19% relative uncertainty),*  
51 *consistent with average literature values and the uncertainties estimates. The*  
52 *uncertainty in EC mass concentration, estimated from measurement repeats based on*  
53 *the EUSAAR2 protocol only, is within 7% in our case. The precision uncertainty in*  
54 *the aethalometer attenuation measurements was estimated as  $15 \text{ Mm}^{-1}$  based on the*  
55 *standard deviation of its signals prior to aerosol being injected into the smog*  
56 *chamber. The MWAA data have an estimated noise level and precision of  $12 \text{ /Mm}$  and*  
57 *10% respectively, and these uncertainties have been added in quadrature to provide*  
58 *the overall uncertainties shown, for example, as error bars in Fig. 1 below. To*  
59 *compare the MWAA and aethalometer measurements, we determined  $b_{\text{abs,MWAA},880\text{nm}}$*   
60 *by extrapolating the absorption coefficients measured at 850 nm to 880 nm using an*  
61  *$\alpha$ -value determined from the ratio between the absorption coefficients at 850 nm and*  
62 *635nm. The uncertainty associated with this extrapolation is considered negligible*  
63 *relative to the overall MWAA uncertainty.*

64

65 → **Possible unquantified uncertainties:**

66 *There are significant uncertainties in the measurement of aerosol absorption using*  
67 *filter-based techniques (e.g., Collaud Coen et al., 2010). Here, we have used MWAA*  
68 *measurements as a reference to scale the aethalometer data, using a single C value.*  
69 *The correction factor C, which accounts for scattering effects within the filter matrix*  
70 *(Drinovec et al., 2015), may depend on the aerosol sample (Collaud Coen et al.,*  
71 *2010). In this study, we evaluated the variability in this factor for our primary and*  
72 *aged samples, by directly comparing the aethalometer to MWAA measurements, as*  
73 *discussed below. The MWAA has been previously validated against a polar*  
74 *nephelometer and a MAAP (Massabo et al., 2013), which, in turn, has been validated*  
75 *against numerous in situ methods (e.g., Slowik et al., 2007). The excellent correlation*  
76 *between MWAA and EC in our study (discussed below) supports the high confidence*  
77 *in the MWAA filter based absorption measurements conducted here. Another*  
78 *significant source of uncertainty in filter-based absorption measurements is the*

79 possible sorption (or evaporation) of volatile organics on (or from) the filter material.  
80 This may lead to an overestimation (or underestimation) of OA absorption. However,  
81 we have minimized sorption artefacts by utilizing a charcoal denuder. We have  
82 obtained an excellent correlation between OA absorption measurements derived from  
83 the MWAA-calibrated aethalometer and from quartz filter samples (see discussion  
84 below, Fig. 6 in the main text and S13 in the supplementary information). Although  
85 both of these techniques involved filter sampling, their sampling timescale is an order  
86 of magnitude different, and a difference is therefore expected if sorption (or  
87 evaporation) caused a substantial bias in our results. We therefore conclude that it is  
88 unlikely that artifacts associated with filter sampling have biased the absorption  
89 measurements. Finally, uncertainties related to pyrolysis during thermo-optical  
90 analysis may bias EC measurements. Such uncertainties arise from unstable organic  
91 compounds, and can be significant for biomass-burning samples, leading to biases on  
92 the order of 20% for EC (e.g. Schauer et al., 2003; Yang and Yu., 2007). To minimize  
93 these biases we applied the EUSAAR2 protocol. The optical properties of such  
94 organics are generally different from BC; therefore, the excellent correlation between  
95 MWAA and EC data in Fig. 1A suggest that pyrolysis effects were not a major source  
96 of uncertainty in our data set.

97 We have added the following abbreviation to the corrected text:

98 L146. multi-angle absorption photometer (MAAP, Petzold and Schönlinner, 2004).

99 We have also expanded the text in some places to reflect the considerations presented above,  
100 as shown in the response to the next comment.

101 ***New references added:***

102 Schauer, J. J., Mader, B. T., Deminter, J. T., Heidemann, G., Bae, M. S., Seinfeld, J.  
103 H., Flagan, R. C., Cary, R. A., Smith, D., Huebert, B. J., Bertram, T., Howell, S.,  
104 Kline, J. T., Quinn, P., Bates, T., Turpin, B., Lim, H. J., Yu, J. Z., Yang, H., and  
105 Keywood, M. D.: ACE-Asia intercomparison of a thermaloptical method for the  
106 determination of particle-phase organic and elemental carbon, *Environ. Sci.*  
107 *Technol.*, 37, 993–1001, <https://doi.org/10.1021/es020622f>, 2003.

108  
109 Slowik, J. G., E. S. Cross, J.-H. Han, P. Davidovits, T. B. Onasch, J. T. Jayne, L. R.  
110 Williams, M. R. Canagaratna, D. R. Worsnop, R. K. Chakrabarty, H. Moosmüller, W.  
111 P. Arnott, J. P. Schwarz, R. S. Gao, D. W. Fahey, G. L. Kok and A. Petzold, An inter-  
112 comparison of instruments measuring black carbon content of soot particles, *Aerosol*  
113 *Sci. Technol.* 41, 3, 295-314, 2007.

114  
115 Collaud Coen, M., Weingartner, E., Apituley, A., Ceburnis, D., Fierz-Schmidhauser,  
116 R., Flentje, H., Henzing, J. S., Jennings, S. G., Moerman, M., Petzold, A., Schmid, O.,  
117 and Baltensperger, U.: Minimizing light absorption measurement artifacts of the  
118 Aethalometer: evaluation of five correction algorithms, *Atmos. Meas. Tech.*, 3, 457-  
119 474, <https://doi.org/10.5194/amt-3-457-2010>, 2010.

120

121

122 *Yang, H. and Yu, J. Z.: Uncertainties in charring correction in the analysis of*  
123 *elemental and organic carbon in atmospheric particles by thermal/optical methods,*  
124 *Environ. Sci. Technol., 36 (23), 5199–5204, 2002.*  
125

126 - In this study the BC particles were not observed to be coated with other kind of particles  
127 (i.e., no internal mixing but external mixing was observed). However, ambient studies have  
128 shown that BC particles in the atmosphere are usually coated and this coating causes an  
129 enhancement of BC absorption. Given that, how representative is this study of “atmospheric  
130 aging”? Were the experiments not long enough to “age” the BC particles?

131 The reviewer raised two points, to which we reply separately.

132 The reviewer questions how representative our study was of atmospheric aging. The  
133 emissions studied here are representative of flaming wood in stoves typically used in Western  
134 Europe, while aging is equivalent to ~2 days of OH-driven photochemistry, under  
135 atmospheric winter day time conditions in the mid-latitude.

136 The reviewer questions the reasons for the lack of lensing with aging, when such effect had  
137 been observed in the field. We have clarified our language here. Our aim has been to assert  
138 that our measurements are poorly represented by a pure core-shell conceptual model of  
139 internal mixing. This assertion is based on our measured absorption coefficients, and we have  
140 modified the text to explain this in more detail as quoted immediately below.

141 The AMS measurements showed that the amount of OA generated during aging was  
142 substantial. Likewise, the SMPS showed a considerable growth of the primary particles with  
143 aging. If BC and OA are naively treated as core-shell mixtures, an absorption enhancement of  
144 ~1.8 would have been predicted, with an average increase in the coating mass by a factor of 3  
145 (see Bond et al., 2006). However, our absorption-coefficient measurements in Fig. 1a showed  
146 that we did not observe any absorption enhancement. Therefore, we do not conclude that “BC  
147 particles were not coated” but rather than “the particles studied could not be represented by a  
148 core-shell description of coatings that envelop the central BC core”. The particles may be  
149 internally mixed, but of a morphology more complex than core-shell – e.g. off-centered  
150 coatings with complexities due to the aggregated morphology of BC, see e.g. the microscopy  
151 images of biomass-burning particles by China et al. 2013. Modelling or even accurately  
152 describing such morphologies is well beyond the scope of our experimental study. Current  
153 literature reports for the lensing effect are conflicting, showing that absorption enhancements  
154 upon significant BC coating can be less than 5% (Cappa et al., 2012) or as large as 150% (Liu  
155 et al., 2015). Recent experimental work suggests that such discrepancies are related to the  
156 complex black carbon morphology and a core-shell description does not adequately capture  
157 mixed-BC optical properties and may considerably overestimate the observed absorption  
158 values (Liu et al., 2017). As lensing effect was negligible in our case, we have assumed that  
159 the aerosol optically behaves as an external mixture between BC and BrC. We note that while  
160 this assumption is important for estimating the BC absorption, the BrC absorption is not very

161 sensitive to the assumed morphology. Based on both reviewers comments and to avoid  
162 confusion we have modified the Results section 4.1 as follows:

163

#### 164 *Section 4.1.*

165 *We have independently determined the  $MAC_{BC}(880nm)$  and the aethalometer  $C$  values*  
166 *under our conditions, as follows. We determined  $MAC_{BC}(880nm)$  from the regression*  
167 *between the absorption coefficients at 880 nm obtained from the MWAA and the EC*  
168 *mass measured by the Sunset analyzer (Fig. 1A). The slope of this regression may be*  
169 *used to estimate the  $MAC_{BC}(880nm)$ , which we retrieved as  $4.7 \pm 0.3 m^2 g^{-1}$  by an*  
170 *uncertainty-weighted linear least-squares fit . The corresponding intercept was not*  
171 *significantly different from zero ( $-3 \pm 3 /Mm$ ). Our  $MAC_{BC}(880nm)$  is not statistically*  
172 *significantly different from the value recommended by Bond et al., (2006) for*  
173 *externally-mixed BC (extrapolating their  $MAC_{BC}(550nm)$  to 880 nm by assuming  $\alpha=1$*   
174 *provides  $MAC_{BC}(880nm)= 4.7 \pm 0.7 m^2 g^{-1}$ ). The strong correlation between*  
175  *$b_{abs,MWAA,880nm}$  and EC in Fig. 1A shows that  $MAC_{BC}(880nm)$  did not vary with aging*  
176 *during our study (see also Fig. S2-a). It also indicates that measurement artefacts for*  
177 *both instruments were negligible, as the fundamental differences between the two*  
178 *techniques mean that any artefacts are unlikely to be similar between them (charring*  
179 *for EC vs. adsorption artefacts for MWAA). Our absorption coefficient measurements*  
180 *also provide insights into particle mixing state in this study. Since a single MAC*  
181 *adequately described our samples at all levels of aging (Fig. 1A and Fig. S2-a), in*  
182 *spite of a factor of 3.3 average increase in the aerosol mass, our samples cannot be*  
183 *adequately described by a core-shell Mie model. Such a core-shell model would*  
184 *predict an absorption enhancement by a factor of  $\sim 1.8$  (Bond et al., 2006) for the*  
185 *observed OA mass increase with aging, which was not observed in our case. This*  
186 *observation is also supported by the time resolved attenuation measurements at 880*  
187 *nm using the aethalometer (Fig. S3), suggesting that little ( $<10\%$ ) to no increase in*  
188 *the attenuation coefficients upon SOA formation. We emphasize that this conclusion*  
189 *does not indicate that no internal mixing occurred, but rather that the simplified*  
190 *concept of negligible mixing better describes our data than the equally simplified*  
191 *concept of a core-shell description of coatings that completely envelop the central BC*  
192 *core. This may be due to the complex morphology of internally-mixed BC, which has*  
193 *been previously observed for wood burning particles (e.g., China et al., 2013; Liu et*  
194 *al., 2015; Liu et al., 2017). It may also be related to the fact that OA and BC are*  
195 *emitted during separate phases of combustion. OA rich particles are emitted during*  
196 *the pre-flaming pyrolysis stage of combustion, whereas most BC is emitted during*  
197 *flaming combustion (Corbin et al., 2015a, 2015b; Haslett et al., 2018; Heringa et al.,*  
198 *2011). These two stages of combustion may coexist in different regions of the stove,*  
199 *particularly during simulated real-world usage. As lensing effect was negligible in*  
200 *our case, we have assumed that the aerosol optically behaves as an external mix*

201 *between BC and BrC during Mie calculation (see section 3.4). We note that while this*  
202 *assumption is important for estimating the BC absorption, the conclusion drawn*  
203 *about the BrC absorption are not very sensitive to the assumed morphology.*

204 *We determined time-resolved wavelength-dependent absorption coefficients as*  
205 *follows. We used the aethalometer to obtain filter attenuation coefficients with high*  
206 *time resolution, which were then calibrated to obtain absorption coefficients by*  
207 *deriving the factor C (Equation 1) using the MWAA measurements of filter samples. C*  
208 *was obtained from an uncertainty-weighted linear least-squares fit as  $3.0 \pm 0.2$  (Fig.*  
209 *1B); the intercept of the fit was not significantly different from zero, within two*  
210 *standard deviations ( $-17 \pm 14$ ). A very strong correlation could be observed between*  
211 *MWAA and aethalometer (Fig. 1B), implying that C is independent of the type of the*  
212 *aerosol sampled (see also Fig. S2-B). Therefore, we used a single C value to obtain*  
213 *time-resolved wavelength-dependent absorption coefficients from the aethalometer*  
214 *attenuation measurements at the different wavelengths for primary and aged aerosols.*

215 *Note that the manufacturer's default values, which were not applied in our case, are*  
216 *1.57 for C (using TFE-coated glass fiber filters) and  $12.2 \text{ m}^2 \text{ g}^{-1}$  for  $\sigma_{\text{ATN}}$  at 880 nm*  
217 *(Gundel et al., 1984, Drinovec et al., 2015). The C value determined here is larger*  
218 *than the manufacturer-default value for the AE33, resulting in smaller absorption*  
219 *coefficients. However, the calculated  $\sigma_{\text{ATN}}$  at 880 nm ( $13.8 \text{ m}^2 \text{ g}^{-1}$ ), which can be*  
220 *retrieved as the product of the C value and  $\text{MAC}_{\text{BC}}(880\text{nm})$  (Eq. (3)), is similar to the*  
221 *factory-default  $\sigma_{\text{ATN}}$ . Therefore, our calibrated  $M_{\text{eBC}}$  concentrations, calculated from*  
222 *the attenuation coefficients using  $\sigma_{\text{ATN}}$  (Eq. (2)), are similar to the factory-default*  
223  *$M_{\text{eBC}}$ . We note that  $M_{\text{eBC}}$  has not been used for  $\text{MAC}_{\text{OA}}$  calculations, and is only used*  
224 *for the calculation of the mass fractions of BC and OA for display purposes (Fig. 2, 3,*  
225 *7 and 8).*

## 226 227 **References:**

228  
229 *China, S., Mazzoleni, C., Gorkowski, K., Aiken, A. C., and Dubey, M. K.: Morphology*  
230 *and mixing state of individual freshly emitted wildfire carbonaceous particles, *Nat.**  
231 **Commun.*, 4, 1–7, doi:10.1038/ncomms-3122-2013, 2013.*

232 *Liu, S., Aiken, A. C., Gorkowski, K., Dubey, M. K., Cappa, C. D., Williams, L. R.,*  
233 *Herndon, S. C., Massoli, P., Fortner, E. C., Chhabra, P. S., Brooks, W. A., Onasch, T.*  
234 *B., Jayne, J. T., Worsnop, D. R., China, S., Sharma, N., Mazzoleni, C., Xu, L., Ng, N.*  
235 *L., Liu, D., Allan, J. D., Lee, J. D., Fleming, Z. L., Mohr, C., Zotter, P., Szidat, S. and*  
236 *Prévôt, A. S. H.: Enhanced light absorption by mixed source black and brown carbon*  
237 *particles in UK winter, *Nat. Commun.*, 8435, doi:10.1038/ncomms9435, 2015b.*

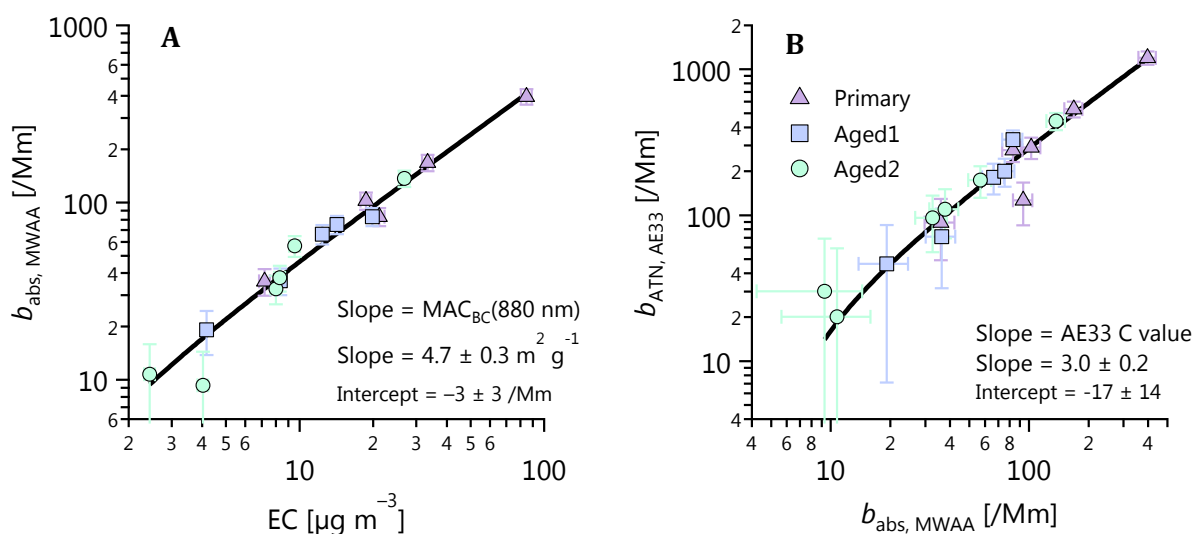
238 *Liu, D. T., Whitehead, J., Alfarra, M. R., Reyes-Villegas, E., Spracklen, D. V.,*  
239 *Reddington, C. L., Kong, S. F., Williams, P. I., Ting, Y. C., Haslett, S., Taylor, J. W.,*  
240 *Flynn, M. J., Morgan, W. T., McFiggans, G., Coe, H., and Allan, J. D.: Black-carbon*

241  
242  
243  
244  
245  
246  
247  
248

absorption enhancement in the atmosphere determined by particle mixing state, *Nat. Geosci.*, 10, 184–188, doi:10.1038/ngeo2901, 2017.

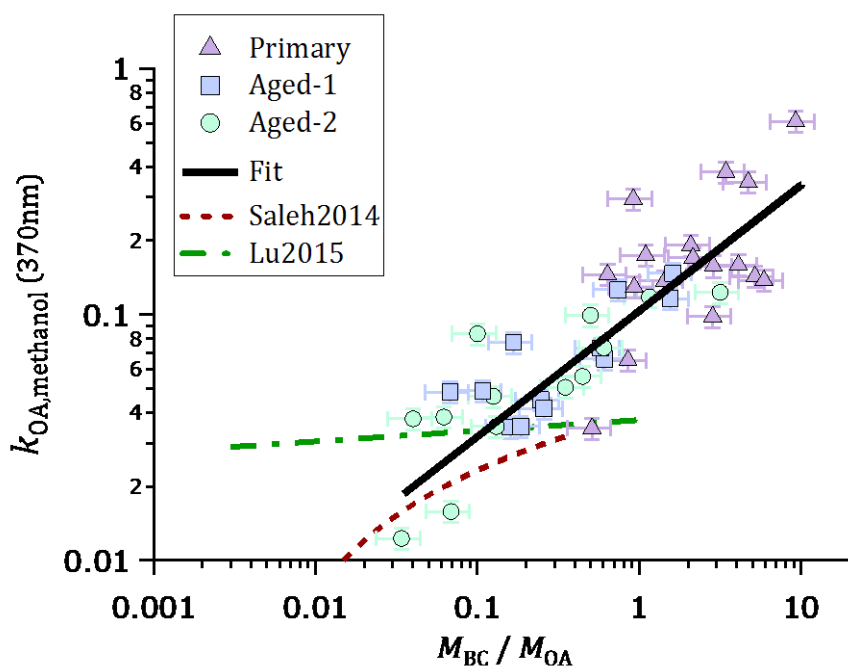
Haslett, S. L.; Thomas, J. C.; Morgan, W. T.; Hadden, R.; Liu, D.; Allan, J. D.; Williams, P. I.; Keita, S.; Lioussse, C. and Coe, H. Highly controlled, reproducible measurements of aerosol emissions from combustion of a common African biofuel source, *Atmos. Chem. Phys.*, 385-403, doi:10.5194/acp-18-385-2018, 2018.

249 We slightly modified Fig. 1A to reflect these changes. We included error bars in Fig. 1 which  
250 had been missing previously.



251  
252

We also included error bars in Fig. 7:



253  
254

255

256 SPECIFIC COMMENTS

257 1. Lines 107-108: How can you guarantee the correction factor C is wavelength  
258 independent?

259

260 This is a good question which has often been neglected in the literature. Recently, Corbin et  
261 al. (2018), in their Section S3.2, presented a very detailed discussion and reanalysis of the  
262 wavelength dependence of the C-value.

263

264 In that publication, the authors described the wavelength dependence of the C-value as  
265 separated into a filter dependence and a scattering cross-sensitivity measurement, and  
266 presented four different arguments for its wavelength dependence being negligible:

267

- 268 1. They compared their wavelength-dependent absorption coefficient measurements with  
269 MWAA measurements (which do not rely on a C-value-like correction) and found good  
270 agreement between the two techniques. We have also verified our AE33 data with  
271 MWAA data in the present study.
- 272 2. They pointed out that the measured AAE would be biased by a wavelength-dependent C  
273 value, so that their measurements of an AAE of 1.0 for samples dominated by BC (in  
274 agreement with extensive literature) indicates a negligible wavelength dependence of C.
- 275 3. They combined measurements of aerosol SSA (ranging from 0.5 to 0.9) with size-  
276 dependent scattering cross-sensitivity measurements to quantitatively estimate the  
277 influence of scattering cross-sensitivity as negligible. SSA measurements were not  
278 available in our study, but our measured size distributions indicate that our particles were  
279 generally small enough that their conclusions can be extrapolated to our samples.
- 280 4. They described previous work where different filter materials were compared, with no  
281 significant effect on the wavelength dependence of the C value (Drinovec et al., 2015).

282

283 In conclusion, the C value is known to depend on the filter material but its wavelength  
284 dependence has been shown to be negligible for samples such as those studied in the  
285 manuscript presently under review.

286

287 We have updated the manuscript as follow:

288

289 *L111. As discussed in detail by Corbin et al. (2018), the wavelength-dependence of C*  
290 *can be expected to be negligible*

291

292

293

294

295

296



297

298 **References:**

299 *Corbin, J. C.; Pieber, S. M.; Czech, H.; Zanatta, M.; Jakobi, G.; Massabò, D.;*  
300 *Orasche, J.; El Haddad, I.; Mensah, A. A.; Stengel, B.; Drinovec, L.; Mocnik, G.;*  
301 *Zimmermann, R.; Prévôt, A. S. H. and Gysel, M. Brown and black carbon emitted by*  
302 *a marine engine operated on heavy fuel oil and distillate fuels: optical properties, size*  
303 *distributions and emission factors, J. Geophys. Res. Atmos., 123, 6175-6195,*  
304 *doi:10.1029/2017JD027818, 2018.*

305 *Drinovec, L., Močnik, G., Zotter, P., Prévôt, A. S. H., Ruckstuhl, C., Coz, E.,*  
306 *Rupakheti, M., Sciare, J., Müller, T., Wiedensohler, A. and Hansen, A. D. A.:*  
307 *The “dual-spot” aethalometer: an improved measurement of aerosol black carbon*  
308 *with real-time loading compensation, Atmos. Meas. Tech., 8(5), 1965–1979,*  
309 *doi:10.5194/amt-8-1965-2015, 2015.*

310 2. Line 122: “1.57 for C”. Does this mean your Aethalometer was using TFE-coated  
311 glass fiber filters? Please mention the filter material.

312

313 Our Aethalometer was using TFE-coated glass fiber filters. We have added some  
314 clarifications in the text as follows:

315

316 *The manufacturer default values are 1.57 for C (TFE-coated glass fiber filters) and*  
317 *12.2 m<sup>2</sup> g<sup>-1</sup> for  $\sigma_{ATN}$  at 880 nm, which corresponds to a  $MAC_{BC}(880\text{ nm})$  of 7.77 m<sup>2</sup> g<sup>-1</sup>*  
318 *(Gundel et al., 1984 , Drinovec et al., 2015).*

319

320 3. Line 128: It should be necessary to add some more discussion about possible artifacts  
321 that affect both techniques (MWAA and Sunset analyser).

322

323 We have added text, cited and quoted in the preceding responses, to better explain how the  
324 MWAA functions and has been validated. We have also added a discussion of artifacts  
325 affecting the Sunset analyzer in the new subsection quoted above. The relevant text  
326 (Section 4.1, as cited at L128) has been updated following this comment and was quoted in  
327 our response to this reviewer’s overall comments. For clarity, we will not quote the text again  
328 here.

329

330 **References:**

331 *Cavalli, F., Viana, M., Yttri, K. E., Genberg, J. and Putaud, J.-P.: Toward a*  
332 *standardised thermal-optical protocol for measuring atmospheric organic and*  
333 *elemental carbon: the EUSAAR protocol, Atmos. Meas. Tech., 3(1), 79–89,*  
334 *doi:doi:10.5194/amt-3-79-2010, 2010.*

335

336

337

338 *Massabò, D., Bernardoni, V., Bove, M. C., Brunengo, A., Cuccia, E., Piazzalunga, A.,*  
339 *Prati, P., Valli, G. and Vecchi, R.: A multi-wavelength optical set-up for the*  
340 *characterization of carbonaceous particulate matter, J. Aerosol Sci., 60, 34–46,*  
341 *doi:10.1016/j.jaerosci.2013.02.006, 2013.*

342

343 4. Line 136. MWAA measurements. Which artifacts are to be considered when using  
344 this technique in comparison to the Aethalometer? Can you provide more evidence on the  
345 comparison of this technique to other absorption measurements like MAAP or PAS?

346

347 This good comment has been addressed in our response to this reviewer's first major  
348 comment. The MWAA has been validated against the MAAP by Massabo et al. (2013).

349 5. Line 171: What do you mean with the online  $k_{OA}$ ? How was this measured?

350 We apologize for this confusion; we have not measured the  $k_{OA}$  online. We corrected this  
351 error and rephrased the sentence for clarity:

352 *If the OA extraction efficiency was less than unity, then the absorption (or MAC)*  
353 *predicted from our solvent-extraction measurements would be less than that measured*  
354 *(or calculated) using our real-time measurements (MWAA-calibrated aethalometer).*

355

356 6. Lines 403-411: How does this result compares to other studies? Ambient  
357 measurements have shown quick oxidation of brown carbon chromophores. Please comment  
358 about it.

359

360 The results of Forrister et al., 2015 suggest that BrC emitted from wildfires has very poor  
361 atmospheric stability with most of the aerosol being lost within a day, either due to chemical  
362 loss or evaporation. However, other ambient and chamber studies clearly (Zhong et al., 2014,  
363 Lee et al., 2014) imply that the optical properties of BrC are strongly dependent on the  
364 complex process of aging. Photo-chemical destruction and formation of chromophores is a  
365 topic of current research and may strongly depend on the gas phase composition.

366

#### 367 **References:**

368 Forrister, H., Liu, J., Scheuer, E., Dibb, J., Ziemba, L., Thornhill, L. K., Anderson, B.,  
369 Diskin, G., Perring, A. E., Schwarz, J. P., Campuzan-Jost, P., Day, D. A., Palm, B. B.,  
370 Jimenez, J. L., Nenes, A., Weber, R. J.: Evolution of brown carbon in wildfire plumes,  
371 *Gophys. Res. Lett.*,42, 4623-4630, doi: 10.1002/2015GL063897, 2015.

372

373

374

375  
376  
377  
378  
379  
380  
381  
382  
383  
384  
385  
386  
387  
388  
389  
390  
391  
392  
393  
394  
395  
396  
397  
398  
399  
400  
401  
402  
403  
404  
405  
406  
407  
408  
409  
410  
411  
412  
413  
414  
415

Zhong, M. and Jang, M.: Dynamic light absorption of biomass-burning organic carbon photochemically aged under natural sunlight, *Atmos. Chem. Phys.*, 14, 1517-1525, doi:10.5194/acp-14-1517-2014, 2014.

Lee, H. J., Aiona, P. K., Laskin, A., Laskin, J. and Nizkorodov, S. A.: Effects of solar radiation on the optical properties and molecular composition of laboratory proxies of atmospheric brown carbon, *Environ. Sci. Technol. Lett.*, 48(17), 10217-10226, doi:10.1021/es502515r, 2014.

Technical Corrections:

7. Line 101: It would be convenient to add numbers to these headlines across the manuscript for the sake of readability (e.g., “2.2.1 Aethalometer”).

We agree, however, as our manuscript still needs to be converted from Word to LaTeX during typesetting, we will leave this decision to the journal staff.

8. Line 123: at?

The ‘at’ was removed from the sentence in line 123.

9. Line 141: Which angles?

We modified the text as follows:

*L140. The first photodiode is placed behind the filter for transmittance measurements (0° relative to the incident light, 1.5 cm from the sample), while the other two photodiodes are positioned at 125° and 165° (11 cm from the sample) to collect the back scattered light.*

10. Lines 294-295: Please rephrase this sentence to improve understanding.

We have rewritten the entire section for improved clarity, as quoted at the beginning of this response.

416  
417  
418  
419  
420  
421  
422  
423  
424  
425  
426  
427  
428  
429  
430  
431  
432  
433  
434  
435  
436  
437  
438  
439  
440  
441  
442  
443  
444  
445  
446  
447  
448  
449  
450  
451  
452  
453  
454  
455  
456  
457  
458

11. Lines 296-299: It can be found awkward that the two variables needed to calculate MAC are coming from the same measurement technique (Aethalometer). Please try to sustain the reasons why it was done this way.

This was not the case, the two variables for the  $MAC_{BC}$  came from the MWA and EC. The aethalometer provided high time resolution attenuation coefficients which could be calibrated to give absorption coefficients needed for the  $MAC_{OA}$  retrieval. eBC was not needed for  $MAC_{OA}$  calculations as now clarified in the text, but for display purposes (Fig. 2, 3, 7 and 8). In the revised manuscript we have clarified the text (see new Section 4.1 quoted at the beginning of this response).

12. Line 372: It should be written “Eq. (19)”, and “Fig. 4”. Please implement this across the manuscript. Check the journal guidelines.

This has been done in the corrected version of the manuscript.

13. Lines 542-543: Please add uncertainty intervals to the reported MAC values.

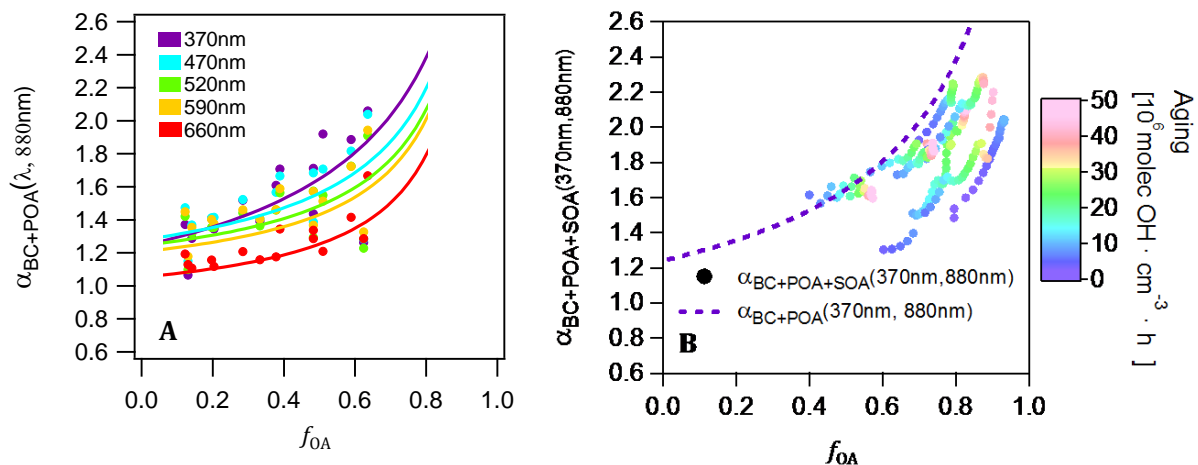
This has been done in the corrected version of the manuscript:

*The mean  $MAC_{SOA}(370nm)$  was  $2.2 m^2 g^{-1}$  (one-sigma variability:  $1.6 - 3.1 m^2 g^{-1}$  according to a  $GSD = 1.39$ ) under our experimental conditions, 2.3 times lower than the mean  $MAC_{POA}(370nm)$  but approximately an order of magnitude higher than MAC values estimated for ambient oxygenated aerosols or reported for SOA from biogenic and traditional anthropogenic precursors.*

459  
 460  
 461  
 462  
 463  
 464

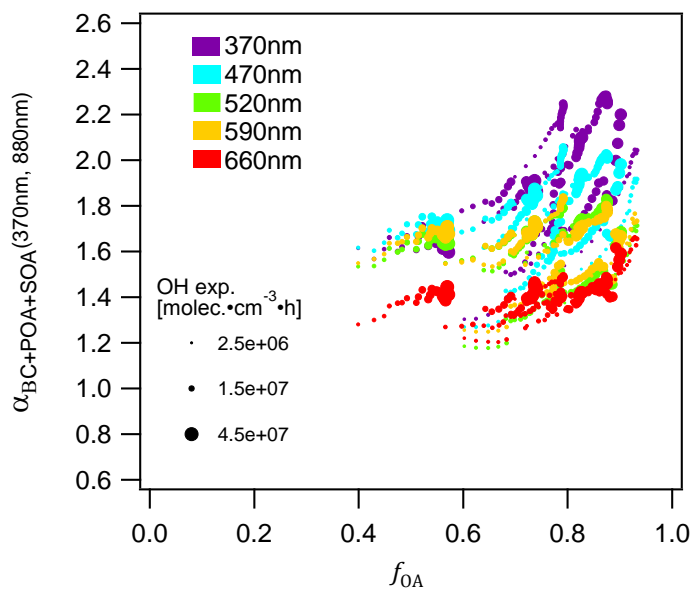
14. Figures 3A, S6, and S10: The data is presented using discrete colors for each wavelength. Please make the legend discrete too.

We have made this change.



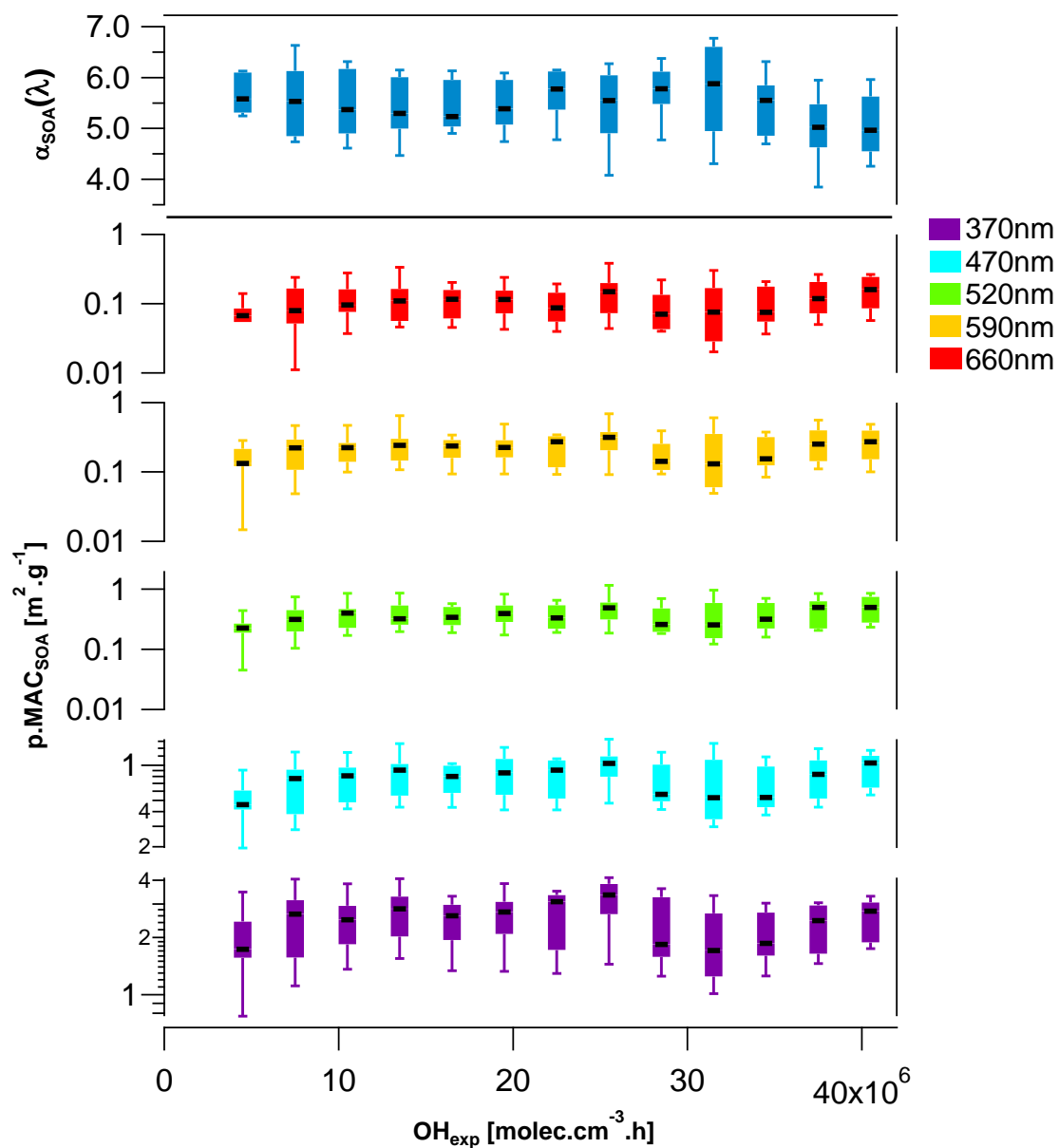
465  
 466

Fig 3A



467  
 468  
 469  
 470

Fig S6



**Fig S10**

15. Figure 5: What do these boxes and whiskers mean? Please clarify.

We added the following to the caption:

*The box marks the 25<sup>th</sup> and 75<sup>th</sup> percentile, while the whiskers mark the 10<sup>th</sup> and the 90<sup>th</sup> percentile.*

486  
487  
488  
489  
490  
491  
492  
493  
494  
495  
496  
497  
498  
499  
500  
501  
502  
503  
504  
505  
506  
507  
508  
509  
510  
511  
512  
513  
514  
515  
516  
517

16. Figure 6: Could you please add the correlation coefficients to the figure?

We are not sure if the reviewer is referring to a Pearson correlation coefficient  $r$  or to the coefficient of determination ( $R^2$ ) of the fit. Although these two quantities are numerically related by a square root function,  $r$  is used to describe a data set and  $R^2$  is used to describe a fit. But  $R^2$  is not appropriate to define our fit, as we performed an orthogonal regression. We have reported the standard error of the fit parameters as appropriate diagnostic statistics. We have followed the reviewer's request to include a correlation coefficient in the paper, but, to avoid confusing the reader, we have mentioned the Pearson correlation coefficient in the text but not added it to the figure, as we expect most readers to misinterpret it as reflecting goodness of fit:

Line 421. *Fig. 6B shows excellent correlation between the  $MAC_{OA}(370nm)$  values obtained from the  $kOA$  of the solvent-extracted OA with the in-situ method described above. The Pearson correlation coefficient was 0.8, for both solvents.*

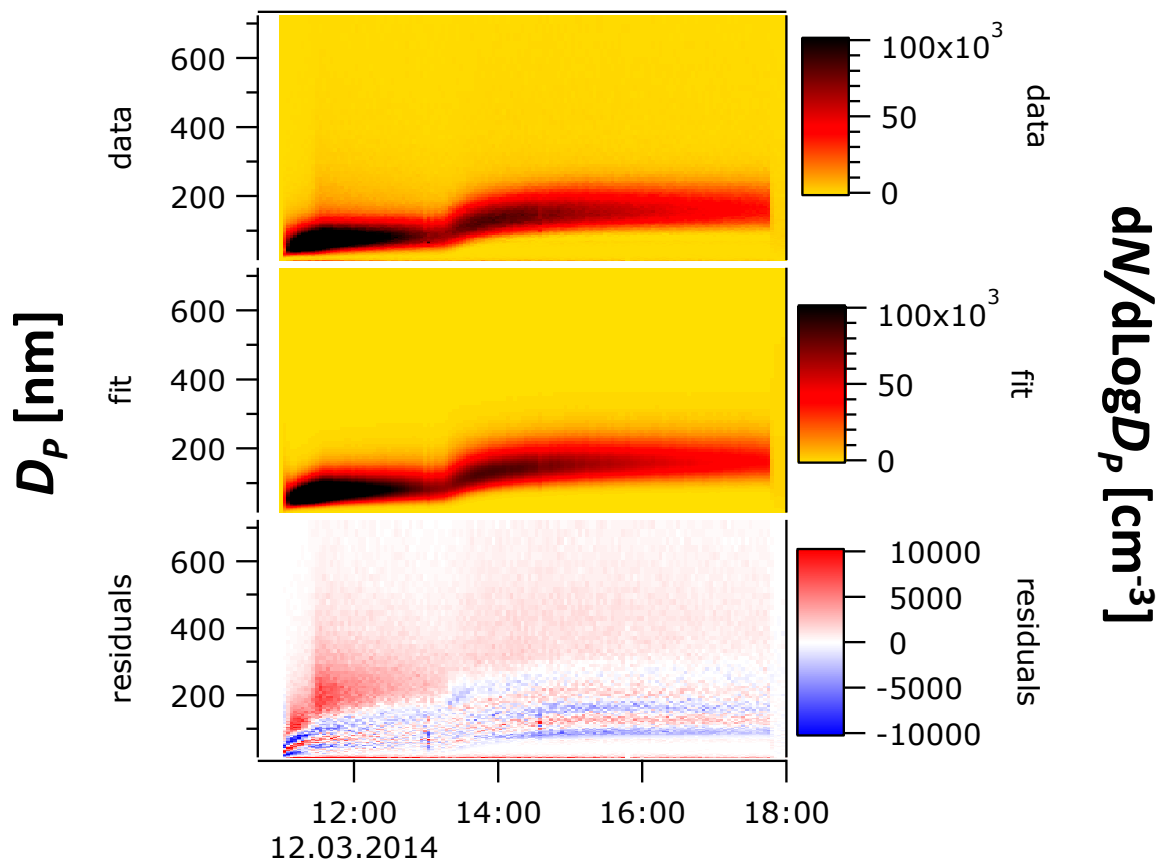
17. Figure S6: I guess you meant  $\alpha$  as a function of  $\lambda$  or do you mean only the wavelength pair 370-880 nm?

We apologize for the confusion; we modified the figure caption as follows:

*Figure S6: Relationship of  $\alpha_{BC+POA+SOA}(\lambda, 880nm)$  to  $f_{OA}$  for seven wavelengths, with symbol sizes indicating OH exposure.*

18. Figures S4, S7, and S9: Please stick to journal guidelines and avoid the use of the jet (or rainbow) color map: "For maps and charts, please keep colour blindness in mind and avoid the parallel usage of green and red. For a list of colour scales that are illegible to a significant number of readers, please visit ColorBrewer 2.0".

Based on the referee comment, we have changed the three figures as follows:

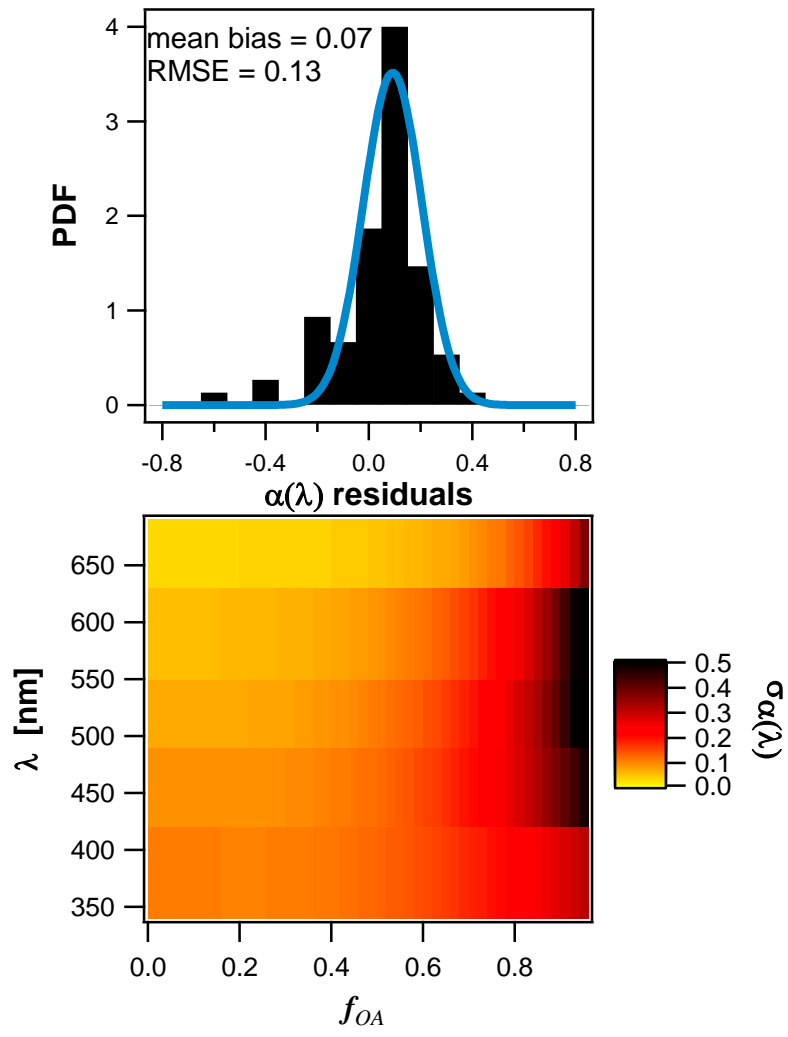


**Fig. S4**

518

519



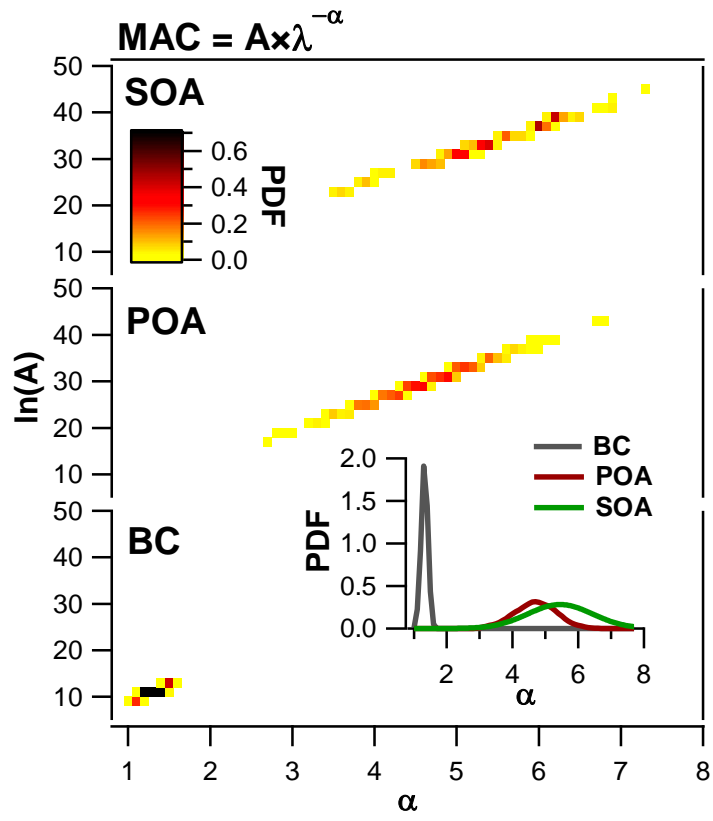


520

521

522

**Fig. S7**



**Fig. S9**

523

524

525

526

527

528

529

530

531

532

533

534 **Anonymous referee 2.**

535 **GENERAL COMMENT**

536 Overall, I find that the authors have presented novel and interesting results on the influence of  
537 photochemical aging on absorption by residential wood combustion emissions. They have  
538 done a better job than I often see in assessing the performance of the aethelometer for their  
539 specific situation, but should include additional details regarding measurement uncertainties,  
540 and how these measurement uncertainties propagate to their final atmospheric implications. I  
541 have numerous comments, mostly just asking for clarification. I believe this paper should be  
542 publishable, with revisions.

543

544 We thank the referee for her/ his constructive comments, which we address below.

545

546 1. Title: It would be good to state “residential wood-burning emissions.”

547

548 The title has been modified in the corrected version of the manuscript:

549

550 *Production of particulate brown carbon during atmospheric aging of residential*  
551 *wood-burning emissions*

552

553 2. L14: It might be good to indicate this was under (likely) high NO<sub>x</sub> conditions.

554

555 The experiments were conducted at estimated NO<sub>x</sub>/NMOG ratios of ~ 0.035 – 0.35 ppm ppm  
556 C<sup>-1</sup> (Bruns et al., 2016). These conditions can be considered as high NO<sub>x</sub>, where most of the  
557 RO<sub>2</sub> radicals react with NO, rather with RO<sub>2</sub>/HO<sub>2</sub>. This information has been added to the  
558 corrected version of the manuscript in section 2.1.

559

560 *Section 2.1.*

561 *Laboratory measurements were conducted in an 8 m<sup>3</sup> Teflon smog chamber (Bruns et*  
562 *al., 2015; Platt et al., 2013) installed within a temperature-controlled housing.*  
563 *Conditions in the chamber were maintained to represent winter time in Europe, i.e.*  
564 *relative humidity ranging between 50 – 90%, at 263 K (Bruns et al., 2015, 2016).*  
565 *Beech wood was combusted in a residential wood stove. Primary emissions were*  
566 *sampled through heated lines at 413 K, diluted by a factor of ~14 using an ejector*  
567 *diluter (DI-1000, Dekati Ltd.), then sampled into the chamber, which provided an*  
568 *additional ten-fold dilution. The overall dilution was a factor of 100 to 200. As we*  
569 *aimed to sample only flaming-phase emissions into the chamber, samples were taken*  
570 *when the modified combustion efficiency (ratio of CO<sub>2</sub> to the sum of CO and CO<sub>2</sub>)*  
571 *was > 0.90. Despite maintaining the same combustion conditions, the resulting*  
572 *organic fraction to the total carbonaceous aerosols in the different samples was*

573 highly variable, indicating that these samples are representative of a mixture of pre-  
574 ignition and flaming emissions (with varying contributions of each combustion stage).  
575 Finally, the resulting NO<sub>x</sub>/NMOG ratios, which dramatically influence SOA  
576 formation through influencing the fate of peroxy radicals, RO<sub>2</sub>, were estimated to be  
577 between 0.035 – 0.35 ppm ppm C<sup>-1</sup> (Bruns et al., 2016). These conditions can be  
578 considered as high NO<sub>x</sub> representative of urban/sub-urban conditions, where most of  
579 the RO<sub>2</sub> radicals react with NO, rather with RO<sub>2</sub>/HO<sub>2</sub>.

580

581 3. L15: It would be good to clarify what wavelength, or if this is integrated in some way.

582

583 This has been clarified in the corrected version of the manuscript:

584

585 *At shorter wavelengths (370 – 470nm), light absorption by brown carbon from*  
586 *primary organic aerosol (POA) and secondary organic aerosol (SOA) formed during*  
587 *aging was around 10 % and 20 %, respectively, of the total aerosol absorption (BrC*  
588 *plus BC).*

589

590 4. L17: Reporting the GSD is informative, but an actual uncertainty estimate would be  
591 better.

592

593 Based on this comment and others below, the reviewer is asking for providing uncertainty  
594 propagation based on measurement precision and accuracy, instead of a GSD which  
595 represents the experimental variability. As we discussed in the reply to the first reviewer, we  
596 have added in the corrected version of the manuscript a new section in the Method,  
597 discussing quantifiable and unquantifiable uncertainties. However, in the result section we  
598 still present the variability in the parameters determined as GSD, as these represent our  
599 confidence levels in these average parameters. We believe that such information is relevant if  
600 these parameters were to be used for future predictions. We also consider that for most of the  
601 parameters experimental variability is much more important than measurement uncertainties  
602 and biases.

603

604 5. L54: It would be good to also cite the work from the Georgia Tech group (Forrister et  
605 al., 2015).

606

607 As suggested by the reviewer the work of Forrister et al., 2015 was cited.

608

609 *Forrister, H., Liu, J., Scheuer, E., Dibb, J., Ziemba, L., Thornhill, L. K., Anderson, B.,*  
610 *Diskin, G., Perring, A. E., Schwarz, J. P., Campuzan-Jost, P., Day, D. A., Palm, B. B.,*  
611 *Jimenez, J. L., Nenes, A., Weber, R. J.: Evolution of brown carbon in wildfire plumes,*  
612 *Gephys. Res. Lett.,42, 4623-4630, doi: 10.1002/2015GL063897, 2015.*

613

614

615

616  
617  
618

6. L229: It is not clear to me why the MAC for POA would be unaffected. The authors write for Eqn. 13 what amounts to:

$$\begin{aligned}\alpha(\lambda, 880) &= \frac{-1}{\ln\left(\frac{\lambda}{880}\right)} \ln\left(\frac{[EC]MAC_{BC\lambda}}{[EC]MAC_{BC880}} + \frac{[OA]MAC_{OA\lambda}}{[EC]MAC_{BC880}}\right) \\ &= \frac{-1}{\ln\left(\frac{\lambda}{880}\right)} \ln\left(\frac{[EC]MAC_{BC\lambda}}{[EC]MAC_{BC880}} + \frac{[OA]MAC_{OA\lambda}}{b_{abs,880}}\right)\end{aligned}$$

619  
620

621 and where they have only included the second line, not the first. The  $b_{abs,880}$  value, while  
622 seemingly independent of the EC measurements as the authors have written them, actually do  
623 depend on the EC measurements because everything has been referenced to the EC  
624 measurement. Perhaps I am simply missing something, but I think that the authors statement  
625 must be further justified. Yes, the MAC of OA is fundamentally independent of the MAC of  
626 BC. But I am not certain that these are practically separated to the extent indicated by the  
627 authors. This is the same challenge that all AAE extrapolation methods face in quantitatively  
628 determining OA absorption in the presence of BC. See e.g. the cited Moosmuller paper or  
629 (*Lack and Langridge, 2013*). Also, the statement on L231 is self-evident. Of course the AAE  
630 at a given wavelength depends on the relative contributions of BC and OA.

631  
632  
633  
634

The reviewer raises two separate points. The first is whether the  $MAC_{OA,\lambda}$  directly depends on the EC mass. The second is whether the BC absorption at a given  $\lambda$  affects the estimation of  $MAC_{OA,\lambda}$ . Below, we address these points separately.

635  
636

1. We consider it inaccurate to say that OA MAC has been referenced to EC. This reflects how we have expressed and applied Equation 13. The intermediate steps leading to equation 13 were intentionally omitted, as they tend to mislead the reader. Equation 13 clearly shows that:

639

- 640 • EC mass concentration is not explicitly required.
- 641 • A potential bias in  $MAC_{BC,880nm}$  due to a bias in EC mass would directly affect the resulting  $MAC_{BC,\lambda}$  in a proportional manner, whereas  $MAC_{OA,\lambda}$  would remain completely unaffected.
- 642 • The resulting  $MAC_{OA,\lambda}$  depends on the input parameter  $M_{OA}$ , thus being affected by potential AMS calibration bias.
- 643 • A potential bias in absorption coefficients measured by the MWAA would propagate to a proportional bias in  $MAC_{OA,\lambda}$ , as aethalometer measurements of  $b_{abs}$  are referenced against the MWAA. Such bias in absorption coefficient would also propagate to a proportional bias in  $MAC_{BC,\lambda}$ , which would happen through a corresponding bias in  $MAC_{BC,880}$
- 644 • The resulting  $MAC_{OA,\lambda}$  also depends on the input parameter  $b_{abs,880nm}$ , which is referenced to the MWAA measurement, whereas the EC data are by no means

645  
646  
647  
648  
649  
650  
651  
652

653 blended into the  $b_{\text{abs}}$  data. Such bias in absorption coefficient would also propagate to  
654 a proportional bias in  $\text{MAC}_{\text{BC},\lambda}$ , which would happen through a corresponding bias in  
655  $\text{MAC}_{\text{BC},880\text{nm}}$ .

656  
657 2. We acknowledge that the determination of  $\text{MAC}_{\text{OA},\lambda}$  depends on the estimated  
658 absorption of BC at a given  $\lambda$ , which in turn depends on the estimated  $\text{MAC}_{\text{BC},\lambda}$ . This  
659 is an issue on any multivariate fitting, where the theoretically independent fitted  
660 quantities are not independently determined. We note that such uncertainties are taken  
661 into account by the fitting errors presented as GSDs in the manuscript. The  $\text{MAC}_{\text{BC},\lambda}$   
662 is similar to an extrapolation of the absorption measurements at  $f_{\text{OA}}=0$ . We note our  
663 experiment covered BC rich particles allowing for an accurate determination of  
664  $\text{MAC}_{\text{BC},\lambda}$ .

665

666 **Reference:**

667

668 Saleh, R.; Adams, P. J.; Donahue, N. M. & Robinson, A. L. The interplay between assumed  
669 morphology and the direct radiative effect of light-absorbing organic aerosol. Geophys. Res.  
670 Lett., 43, 8735-8743, doi:10.1002/2016GL069786, 2016.

671

672 We have added the statements below to the manuscript:

673

674 L222. [...] Equation 10. We have intentionally formulated of Equation 13 as such to  
675 highlight that the retrieved  $\text{MAC}_{\text{OA}}(t,\lambda)$  depends mainly on the input  $M_{\text{OA}}$ .  
676 Correspondingly, the retrieved  $\text{MAC}_{\text{OA}}(t,\lambda)$  is mainly sensitive to potential AMS  
677 calibration biases.

678

679 7. L264: What about chemically induced changes of POA mass, as opposed to just  
680 absorption?

681

682 We have mentioned the changes in POA mass and MAC in the sentence before, in L263:  
683 “Equation 19 is based on the assumption that POA is “chemically inert”, i.e. no chemically  
684 induced changes of  $M_{\text{POA}}$  and  $\text{MAC}_{\text{POA}}$  occur”. We have adjusted the text based on the  
685 Reviewer’s comment to avoid confusion:

686

687 *Such chemically induced changes of absorption coefficient by POA, through a*  
688 *change of  $M_{\text{POA}}$  or  $\text{MAC}_{\text{POA}}$ , if they occur, are assigned to the absorption by SOA,*  
689 *thus resulting in a corresponding adjustment of the inferred  $\text{MAC}_{\text{SOA}}$ .*

690

691

692

693

694

695

696

697  
698  
699  
700  
701  
702  
703  
704  
705  
706  
707  
708  
709  
710  
711  
712  
713  
714  
715  
716  
717  
718  
719  
720  
721  
722  
723  
724  
725  
726  
727  
728  
729  
730  
731  
732  
733  
734  
735  
736  
737  
738

8. L287: It is unclear what is meant by “a higher than measured lensing effect.”

This text has been rewritten following the comments above. The revised text clarifies this meaning and, to avoid confusion, we will simply refer to the revised text in response to Reviewer #1’s first comment. The corrected text reads as follows:

#### **Section 4.1.**

*We have independently determined the  $MAC_{BC}(880nm)$  and the aethalometer  $C$  values under our conditions, as follows. We determined  $MAC_{BC}(880nm)$  from the regression between the absorption coefficients at 880 nm obtained from the MWAA and the EC mass measured by the Sunset analyzer (Figure 1A). The slope of this regression may be used to estimate the  $MAC_{BC}(880nm)$ , which we retrieved as  $4.7 \pm 0.3 \text{ m}^2\text{g}^{-1}$  by an uncertainty-weighted linear least-squares fit. The corresponding intercept was not significantly different from zero ( $-3 \pm 3 /Mm$ ). Our  $MAC_{BC}(880nm)$  is not statistically significantly different from the value recommended by Bond et al., (2006) for externally-mixed BC (extrapolating their  $MAC_{BC}(550nm)$  to 880 nm by assuming  $\alpha=1$  provides  $MAC_{BC}(880nm)= 4.7 \pm 0.7 \text{ m}^2\text{g}^{-1}$ ). The strong correlation between  $b_{abs,MWAA,880nm}$  and EC in Fig. 1A shows that  $MAC_{BC}(880nm)$  did not vary with aging during our study (see also Fig. S2-a). It also indicates that measurement artefacts for both instruments were negligible, as the fundamental differences between the two techniques mean that any artefacts are unlikely to be similar between them (charring for EC vs. adsorption artefacts for MWAA). Our absorption coefficient measurements also provide insights into particle mixing state in this study. Since a single MAC adequately described our samples at all levels of aging (Fig. 1A and Fig. S2-a), in spite of a factor of 3.3 average increase in the aerosol mass, our samples cannot be adequately described by a core-shell Mie model. Such a core-shell model would predict an absorption enhancement by a factor of  $\sim 1.8$  (Bond et al., 2006) for the observed OA mass increase with aging, which was not observed in our case. This observation is also supported by the time resolved attenuation measurements at 880 nm using the aethalometer (Fig. S3), suggesting that little ( $<10\%$ ) to no increase in the attenuation coefficients upon SOA formation. We emphasize that this conclusion does not indicate that no internal mixing occurred, but rather that the simplified concept of negligible mixing better describes our data than the equally simplified concept of a core-shell description of coatings that completely envelop the central BC core. This may be due to the complex morphology of internally-mixed BC, which has been previously observed for wood burning particles (e.g., China et al., 2013; Liu et al., 2015; Liu et al., 2017). It may also be related to the fact that OA and BC are emitted during separate phases of combustion. OA rich particles are emitted during the pre-flaming pyrolysis stage of combustion, whereas most BC is emitted during flaming combustion (Corbin et al., 2015a, 2015b; Haslett et al., 2018; Heringa et al., 2011). These two stages of combustion may coexist in different regions of the stove,*

739 particularly during simulated real-world usage. As lensing effect was negligible in  
740 our case, we have assumed that the aerosol optically behaves as an external mix  
741 between BC and BrC during Mie calculation (see section 3.4). We note that while this  
742 assumption is important for estimating the BC absorption, the conclusion drawn  
743 about the BrC absorption are not very sensitive to the assumed morphology.

744 We determined time-resolved wavelength-dependent absorption coefficients as  
745 follows. We used the aethalometer to obtain filter attenuation coefficients with high  
746 time resolution, which were then calibrated to obtain absorption coefficients by  
747 deriving the factor C (Equation 1) using the MWAA measurements of filter samples. C  
748 was obtained from an uncertainty-weighted linear least-squares fit as  $3.0 \pm 0.2$  (Fig.  
749 1B); the intercept of the fit was not significantly different from zero, within two  
750 standard deviations ( $-17 \pm 14$ ). A very strong correlation could be observed between  
751 MWAA and aethalometer (Fig. 1B), implying that C is independent of the type of the  
752 aerosol sampled (see also Fig. S2-B). Therefore, we used a single C value to obtain  
753 time-resolved wavelength-dependent absorption coefficients from the aethalometer  
754 attenuation measurements at the different wavelengths for primary and aged aerosols.

755 Note that the manufacturer's default values, which were not applied in our case, are  
756 1.57 for C (using TFE-coated glass fiber filters) and  $12.2 \text{ m}^2 \text{ g}^{-1}$  for  $\sigma_{\text{ATN}}$  at 880 nm  
757 (Gundel et al., 1984, Drinovec et al., 2015). The C value determined here is larger  
758 than the manufacturer-default value for the AE33, resulting in smaller absorption  
759 coefficients. However, the calculated  $\sigma_{\text{ATN}}$  at 880 nm ( $13.8 \text{ m}^2 \text{ g}^{-1}$  see Fig. 1C), which  
760 can be retrieved from Fig. 1-C or as the product of the C and  $\text{MAC}_{\text{BC}}(880\text{nm})$   
761 (Equation 3), is almost equal to the factory default  $\sigma_{\text{ATN}}$ . Therefore, our  $M_{\text{eBC}}$ , which  
762 can be directly calculated from the attenuation coefficients using  $\sigma_{\text{ATN}}$  (Equation 2),  
763 would be consistent with factory default  $M_{\text{eBC}}$ . We note that  $M_{\text{eBC}}$  has not been used  
764 for MAC calculations. However for displaying purposes, we have calculated time  
765 resolved  $M_{\text{eBC}}$  using  $\sigma_{\text{ATN}}$  and  $b_{\text{ATN}}$  to estimate fraction of BC and OA in the aerosol  
766 (Fig. 2, 3, 7 and 8).

767  
768 9. L290: The authors mention uncertainties here for the UV-Vis measurements. But what  
769 about for all of the in situ or other measurements? This includes [EC], [OA], absorption.  
770 Further discussion of uncertainties is necessary.

771  
772 We thank the reviewer for emphasizing the importance of uncertainties. In the previous  
773 version, uncertainties related with the different measurements were discussed very briefly in  
774 sections related with the description of the measurement techniques. As this comment was  
775 raised by both reviewers, we have added a new section discussing quantifiable and non-  
776 quantifiable uncertainties, which is referenced in the corresponding response and will not be  
777 duplicated here for clarity.



781

782 10. L305: It is not clear how this uncertainty estimate was arrived at. Also, this differs  
783 from the figure. Finally, it is not clear whether this fit has been forced to zero or not.

784

785 The reviewer is right, the uncertainty value in the figure and the text are not the same and the  
786 correct value is the one indicated in the figure (this has been changed accordingly). The  
787 uncertainty is the standard error of the slope of the uncertainty-weighted linear least-squares  
788 fit. The fit has not been forced through zero. Rather, we have fit the full linear model and test  
789 for the significance of the intercept term. The intercept was at  $-3 \pm 3 \text{ m}^2/\text{g}$ , not statistically  
790 significantly different than 0. This intercept was missing from the submitted manuscript and  
791 has now been reported, together with further modifications to the text as quoted in our  
792 modified Section 4.1 above.

793

794 11. L307: Bond et al. did not “report” a value at 880 nm. They reported at a shorter  
795 wavelength. This value is inferred assuming an AAE = 1. It should be noted as such.

796

797 The reviewer is right and the text has been modified accordingly, as specified in the quoted  
798 response to Reviewer #1’s first comment.

799

800 12. L308: While the authors obtain a value of the MAC in good agreement with  
801 extrapolated values from Bond et al., it is not clear to me how this definitely indicates no  
802 lensing effect is present. An uncertainty analysis is necessary. What if, hypothetically, the EC  
803 was biased high and the absorption biased low? The obtained MAC might appear in  
804 agreement with literature, but only within the bounds of the measurement uncertainty. This  
805 statement should be quantified.

806

807 The reviewer raises a valid point that unidentified measurement biases in opposing directions  
808 may have led to the illusion of agreement between our measured  $MAC_{BC}(880\text{nm})$  values and  
809 the literature values for bare BC. The major issue here is that of referencing our absorption  
810 data and EC measurements to reliable and calibrated technique. Absorption measurements  
811 were obtained using MWAA, which has been validated as described in the text, and biases  
812 are expected to be within 10%. Measurement biases related to total carbon measurement are  
813 negligible (within 5%). The high correlation between absorption and EC measurements also  
814 indicates that unquantifiable uncertainties and artefacts (e.g. charring for EC and vapor  
815 adsorption artefacts for MWAA) are negligible, as the fundamental differences between the  
816 two measurements mean that any artefacts are unlikely to be similar between them.  
817 Therefore, measurement biases and unquantifiable artefacts are unlikely to affect the estimated  
818 values for  $MAC_{BC}(880\text{nm})$  and our conclusions about the lack of lensing. We also note that  
819 such conclusions are also supported by independent time resolved attenuation measurements  
820 by the aethalometer, suggesting that little (<10%) to no increase in the attenuation  
821 coefficients upon SOA formation.

822 Our analysis has combined multiple analytical techniques and found good agreement between  
823 all of them. This good agreement reduces the likelihood that opposing measurement biases  
824 led, by chance, to our measurements being in agreement with literature. While it remains

825 theoretically possible that unknown uncertainties and biases existed in our analysis, it is by  
826 Occam's razor more probable that our measurements were in fact accurate and that our cross-  
827 validation by employing different techniques was successful.

828 In the corrected version of the manuscript, we have added a new section discussing the  
829 measurement biases estimated for the different measurements. We have additionally  
830 discussed potential unquantifiable uncertainties. For clarity, we do not duplicate the modified  
831 text here, but we quote the response to reviewer 1.

832

833 13. L315: It is clear that the distribution is reasonably log-normal with a single mode  
834 when considered in number space. But what about in surface area or in volume space, which  
835 is important for the calculation of the MAC\_OA\_bulk (Eqn. 20)? Also, to what extent does  
836 the volume-weighted distribution exceed the SMPS bounds?

837

838 The absolute SMPS volume is not as important for calculating  $MAC_{OA,bulk}$  as the reviewer  
839 understood. As can be seen from Equation 20, we did not use the SMPS data to calculate total  
840 OA volume. We measured total OA mass with an AMS, converted mass into volume using  
841 an assumed density (the assumed density has no impact on the results as we use the same  
842 density to calculate MAC), and then "distributed" this volume across the size distribution  
843 measured by SMPS. That is, the SMPS data provide only a weighting factor for the size  
844 dependence of absorption, which means that uncertainties in these data do not have a major  
845 effect on our results. We performed the calculations in this way to minimize associated  
846 uncertainties, but acknowledge that substantial uncertainties may result. We have estimated  
847 that these may be on the order of 20%, based on the magnitude with which particulate  
848 absorption varies as a function of size (according to Mie theory).

849

850 14. L325 and Fig. S2: I do not understand why in Fig. S2 it says the "reported value" is  
851 2.6 while here 3.0 is given. This should be clarified.

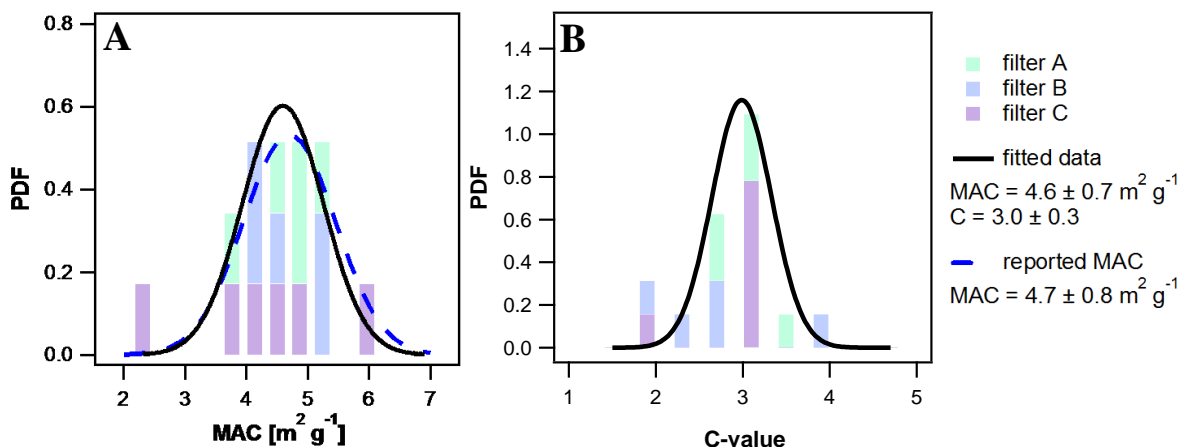
852

853 The reported C value is calculated using a  $\sigma_{ATN} = 12.2 \text{ m}^2 \text{ g}^{-1}$ , as given by the manufacturer,  
854 and  $MAC_{BC}(880\text{nm}) = 4.7 \text{ m}^2 \text{ g}^{-1}$ . The C value = 3 is determined from our attenuation and  
855 absorption measurements which is used in our calculations. This has been discussed in lines  
856 330-336.

857

858 For better clarity, we have now omitted the reported C-value from Fig. S2:

859



860  
 861 *Figure S2: (A) Probability density function (PDF) comparing the MAC values determined by*  
 862 *normalizing MWAA absorption measurements of offline primary (filter A), slightly aged*  
 863 *(filter B: Aged1) and aged (filter C: Aged2) samples to EC (EUSAAR2) measurements of the*  
 864 *same samples (bold line). A literature value for pure BC is also shown (Bond et al., 2006)*  
 865 *(dashed blue line). (B) PDF comparing aethalometer attenuation measurements at 880 nm*  
 866 *and MWAA absorption measurements at 850 nm to retrieve the aethalometer C value.*

867  
 868 15. L332: Not only wood combustion, but really any aerosol if it is assumed that the true  
 869 MAC for BC at 880 nm is  $\sim 4.7 \text{ m}^2/\text{g}$ .

870  
 871 We agree with the reviewer. We specify wood burning because it is the focus of our study.  
 872 This section has been substantially modified and this sentence was removed.

873  
 874 16. L336: This seems circular to me, if the authors are using  $C = 3$  in coming to this  
 875 conclusion. This is demanded through the various inequalities.

876  
 877 We do not think that the math here is circular. Below, we present in bullet points the  
 878 approach followed:

- 879
- 880 •  $\sigma_{\text{ATN}} = \text{MAC}_{\text{BC}} * C$
  - 881 • The Aethalometer provides  $b_{\text{ATN}}$
  - 882 • We determine a C value of 3 from Equation 1 of the paper:
- $$883 \quad C = b_{\text{ATN,AE33}} / b_{\text{abs,MWAA}}$$
- 884 • We determine a  $\text{MAC}_{\text{BC}}$  of  $4.6 \text{ m}^2 \text{g}^{-1}$  by a fit through  $b_{\text{abs,MWAA}}$  and EC thermo-optical  
 885 measurements for primary and aged filter samples.
  - 886 • We have therefore determined the two variables required to calculate  $\sigma_{\text{ATN}}$  and  
 887 determined it as  $13.8 \text{ m}^2 \text{g}^{-1}$ .
  - 888 • We then compare the  $\sigma_{\text{ATN}}$  determined by us to the manufacturer value of  $12.2 \text{ m}^2 \text{g}^{-1}$ .
  - 889 • We then concluded that while the factory default  $b_{\text{abs}}(\lambda)$  has a substantial bias the eBC  
 890 mass determined using default  $\sigma_{\text{ATN}}$  is consistent with the EC mass. We have  
 891 rewritten Section 4.1, as quoted at the beginning of this response, to clarify this  
 892 reasoning.

893  
 894

895  
896  
897  
898  
899  
900  
901  
902  
903  
904  
905  
906  
907  
908  
909  
910  
911  
912  
913  
914  
915  
916  
917  
918  
919  
920  
921  
922  
923  
924  
925  
926  
927  
928  
929  
930  
931  
932  
933  
934  
935  
936  
937  
938

17. Fig. S5 must indicate which studies are being used to define the literature bounds, and also note that this is not the entire range of reported values over the many papers on this topic. This is a subset of values. Consider e.g. (Lewis *et al.*, 2008) or (Liu *et al.*, 2013). The authors have selected a very narrow subset of literature results to present here, and to reference in the text.

While we agree with the reviewer here, one reason that we have not cited e.g. Lewis *et al.* (2008) is that those authors did not experiment on log-wood burning in a modern stove but rather simulated wildfires. Considering the very low emphasis placed on Fig. S5 in our manuscript, and the importance of avoiding confusion between log-wood burning and wildfire burns, we have simply removed the  $\alpha$  ranges in Fig. S5 is only for primary biomass burning emissions. We have also modified the legend to show the two wavelengths of  $\alpha$ .

18. Fig. 3: It is not clear why the data in Fig. A do not overlap with the data in Fig. B. Presumably the data in figure B evolved from the data in Fig. A. Also, in Fig. A it is not clear if each point is for one experiment or whether the variability in  $f_{OA}$  is due to variability within an experiment.

Fig. 3A refers to primary emissions and Fig. 3B refers to secondary emissions (as indicated in the caption), so yes, the presumption here was correct, but only partially. More data are included in Fig. 3A because not all burn experiments were aged. This information has been added to the figure caption. In Fig. 3A each point represents an experiment; therefore the variability in  $f_{OA}$  is due to the variability in emission composition between experiments. Meanwhile, in Fig. 3B the variability in  $f_{OA}$  is due to SOA formation with aging; data from several experiments are shown which explains the wide range of  $f_{OA}$  at low OH exposures. Based on the reviewer comment, we have added the following clarifications to the Fig. 3 caption:

*Figure 3: (A) Relationship of  $\alpha_{BC+POA}(\lambda, 880nm)$  to  $f_{OA}$  for seven wavelengths for primary emissions. Data are colored by the wavelength. Curves are fits of Equation 13 to the data. Each point represents the average of one experiment and therefore the variability in  $f_{OA}$  is related to the variability in the emission composition between experiments. (B) Relationship of  $\alpha_{BC+POA+SOA}(370nm, 880nm)$  to  $f_{OA}$  for several experiments for aged aerosols. Data are color coded by the OH exposure. The variability in  $f_{OA}$  is due to SOA formation with aging; data from several experiments are shown which explains the wide range of  $f_{OA}$  at low OH exposures. Note that more data are included in A than B, as primary emissions for some experiments were not aged.*

939  
940  
941  
942  
943  
944  
945  
946  
947  
948  
949  
950  
951  
952  
953  
954  
955  
956  
957  
958  
959  
960  
961  
962  
963  
964  
965  
966  
967  
968  
969  
970  
971  
972  
973  
974  
975  
976  
977  
978  
979  
980  
981  
982

19. L344: I do not dispute that the AAE values increase with  $f_{\text{OA}}$ . However, it is evident that as wavelength decreases the difference from  $\text{AAE} = 0.9\text{-}1.1$  and the observations increases. This is not clear from the discussion here.

We have stated that the AAE values at low  $f_{\text{POA}}$  are close to those reported for pure BC. We note that for  $\lambda = 660 \text{ nm}$ , the AAE value extrapolated at  $f_{\text{POA}} = 0$ , is equal to 1.04, while for all other wavelengths the value is statistically similar, equal to  $\sim 1.2$ . It can be seen from Fig. 4 that  $\text{MAC}_{\text{BC}}(\lambda)$  profile can be clearly described by a power law, consistent with a constant AAE. To avoid confusion, we have modified the text as follows:

L344. *The  $\alpha(\lambda, 880\text{nm})$  is slightly higher than that of pure BC ( $\sim 1.2$ ; Bond et al., 2013; Zotter et al., 2017) for small  $f_{\text{POA}}$ , while increasing  $f_{\text{POA}}$  corresponded to a distinct increase of  $\alpha(\lambda, 880\text{nm})$ .*

20. L346: The range reported is inconsistent with what is shown in the graph.

The range presented before denoted the P10 and P90; we apologize that we have forgotten to mention this in the text. Based on the reviewer comment and to avoid confusions we have replaced the [P10-P90] by the total range:

*The  $f_{\text{POA}}$  ranges from **0.12 to 0.63**, which is lower than  $f_{\text{POA}}$  reported for open burning emissions (e.g.,  $f_{\text{POA}} \sim 0.75$ , Ulevicius et al (2016)), because our wood-stove emissions feature a more efficient combustion.*

21. L350: this could be strengthened simply by showing a graph of the observations as a function of wavelength, and showing that a single AAE value does not provide for a good fit.

This is shown in Fig. S5 of the SI. We have added in the corrected version of the manuscript a reference to Fig. S5 and modified the figure caption to highlight the point raised by the reviewer as follows:

*Figure S5: Absorption coefficients of fresh wood burning emissions measured using an aethalometer normalized to the eBC mass as a function of wavelength. In the legend each color denotes the  $\alpha_{\text{BC+POA}}(370\text{nm}, 880\text{nm})$  for an individual experiment. The dashed lines mark the absorption profiles calculated using the literature range of  $\alpha$  values obtained for primary biomass burning emissions. The observed absorption spectra have steeper gradients with decreasing wavelength compared to the lines of constant alpha. The systematic decrease in  $\alpha(\lambda, 880\text{nm})$  with increasing  $\lambda$  reflects the more-efficient light absorption by BrC at shorter wavelengths (Moosmüller et al., 2011), and shows that the power law wavelength dependence is an inaccurate oversimplification for this mixed aerosol.*

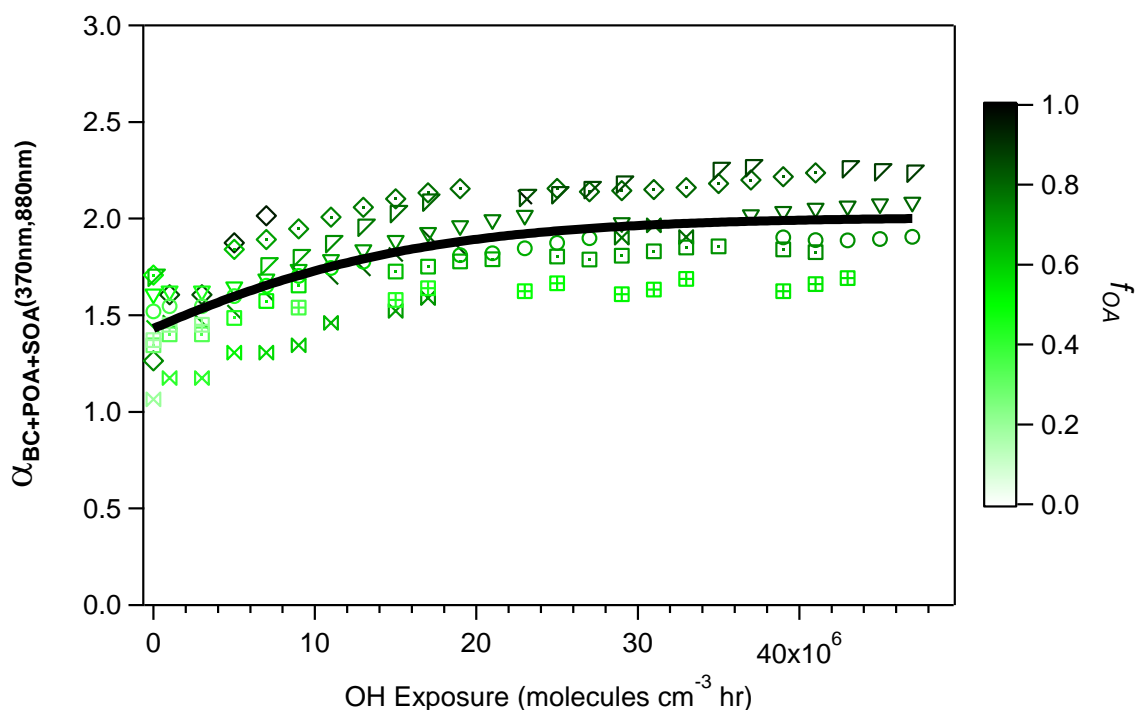
983  
984  
985  
986  
987  
988  
989  
990  
991  
992  
993  
994  
995  
996  
997

We have added the following in the text in L350:

*As illustrated in Fig. S5, the observed absorption spectra have steeper gradients with decreasing wavelength compared to the lines of constant alpha. Such systematic increase in  $\alpha(\lambda, 880\text{nm})$  with decreasing  $\lambda$  reflects the more-efficient light absorption by BrC at shorter wavelengths (Moosmüller et al., 2011), and shows that the power law wavelength dependence is an inaccurate oversimplification for this mixed aerosol.*

22. Fig. 2: For consistency, it would be helpful if the color scale were labeled as  $f_{\text{OA}}$ , similar to Fig. 3 and the text.

This has been modified in the corrected version of the manuscript:



998  
999  
1000  
1001  
1002  
1003  
1004  
1005  
1006  
1007  
1008  
1009  
1010

1011  
1012  
1013  
1014  
1015  
1016  
1017  
1018  
1019  
1020  
1021  
1022  
1023  
1024  
1025  
1026  
1027  
1028  
1029  
1030  
1031  
1032  
1033  
1034  
1035  
1036  
1037  
1038  
1039  
1040  
1041  
1042  
1043  
1044  
1045  
1046  
1047  
1048  
1049  
1050  
1051  
1052  
1053  
1054

23. L356: Looking at Fig. 3B, it is not clear that this is generally the case. The highest  $f_{OA}$  in Fig. 3B does not have the highest AAE. Perhaps the authors mean this just for the high OH exposures. If so, they might consider plotting AAE vs.  $f_{OA}$  for subsets of data binned according to OH exposure. But, as presented, it is not evident that this is a fully general conclusion.

We do mean at higher OH exposures. This has been added in the corrected version of the manuscript:

*Also, note in Fig. 2 that at highest OH exposures, the highest  $\alpha_{BC+POA+SOA}(370nm, 880nm)$  were reached, on average 1.8, during experiments where the  $f_{OA}$  was highest.*

24. L362: While I don't necessarily disagree with this point, I will reiterate that the relationship between Fig. 3A and 3B is not clear. The authors give a dashed curve, but it is not clear how, for example, the data in Fig. 3B that start at such low AAE values at high  $f_{OA}$  values come from Fig. 3A. A stronger connection needs to be made to make this clearer

We think that the misunderstanding comes from the fact that we had not adequately highlighted that not all experiments in Fig. A are shown in Fig. B, as for some of the experiments the emissions were not aged. We chose to represent use all the data available in Fig. A to increase our statistics and expand the  $f_{OA}$  range. As mentioned above this information has been added and the Figure caption now reads as follows:

*Figure 3: (A) Relationship of  $\alpha_{BC+POA}(\lambda, 880nm)$  to  $f_{OA}$  for seven wavelengths for primary emissions. Data are colored by the wavelength. Curves are fits of Equation 13 to the data. Each point represents the average of one experiment and therefore the variability in  $f_{OA}$  is related to the variability in the emission composition between experiments. (B) Relationship of  $\alpha_{BC+POA+SOA}(370nm, 880nm)$  to  $f_{OA}$  for several experiments for aged aerosols. Data are color coded by the OH exposure. The variability in  $f_{OA}$  is due to SOA formation with aging; data from several experiments are shown which explains the wide range of  $f_{OA}$  at low OH exposures. Note that more data are included in A than B, as primary emissions for some experiments were not aged.*

1055  
1056  
1057  
1058  
1059  
1060  
1061  
1062  
1063  
1064  
1065  
1066  
1067  
1068  
1069  
1070  
1071  
1072  
1073  
1074  
1075  
1076  
1077  
1078  
1079  
1080  
1081  
1082  
1083  
1084  
1085  
1086  
1087  
1088

25. L364: For the data in Fig. 3B, extrapolation to  $f_{OA} \rightarrow 0$  for wavelengths  $< 600$  nm suggests an AAE  $\sim 1.2$ - 1.3, larger than the 0.9-1.1 range the authors have taken for BC. This is consistent with the derived  $MAC(370) = 13.7$  m<sup>2</sup>/g for BC, given the value at 880 nm. There is, however, a bit of an inconsistency with how the authors compare with Bond. They state that 13.7 is within the 95% confidence interval of the 11.1 m<sup>2</sup>/g value reported by Bond. But, they have also stated that the AAE = 0.9-1.1. If this is the case, then isn't the range actually narrower? Really, my question here is about the consistency of the statistical interpretation/uncertainty representation.

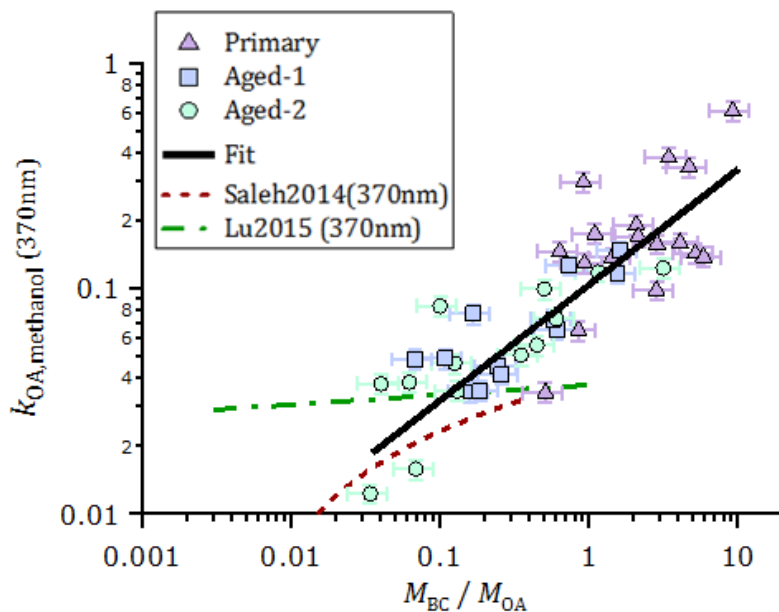
We thank the reviewer once again for raising opportunities for clarifying the text, where some of the information were missing. The range of  $MAC_{BC}(370)$  we calculate is based on an error propagation calculation considering not only the range of AAE reported (0.9-1.1) but also the errors on the absolute  $MAC_{BC}(520)$ . This has now been clarified in the text:

*The obtained fit value of  $MAC_{BC}(370nm)$  was  $13.7$  m<sup>2</sup> g<sup>-1</sup> (GSD 1.1, one-sigma uncertainty 12.4—15.1 m<sup>2</sup>/g), higher but not statistically significantly different from the range estimated based on Bond et al. (2013), considering the uncertainties on both the  $\alpha_{BC}$  values and the  $MAC_{BC}(520nm)$ .*

26. Fig. 7: I find the legend to be incomplete in that it leaves the reader thinking that the Lu and Saleh measurements are from methanol extraction, which they are not. Also, for Saleh (2014), the authors do not at all make clear that the Saleh measurements are at 550 nm, not 370. This is not a fair comparison. Neither is which fuel type of Saleh's has been considered. The authors should provide a fuller picture. Also, the Saleh reference is missing from the bibliography.

Thank you for pointing out the missing reference. The Saleh et al. (2014) data in Fig. 7 were extrapolated to 370 nm using the wavelength dependence of  $k_{OA}$  i.e.  $k_{OA} = k_{OA,550} \times (550 / \lambda)^w$ , given by those authors. Likewise, the Lu et al., 2015 data have also been determined at 370 nm using the wavelength dependence of  $k_{OA}$  provided by the authors ( $k_{OA} = 0.017 \times (550 / \lambda)^{1.62}$ ) and the parameterization of  $k_{OA}$  against BC/OA ratio. We have updated the figure legend and caption according to the reviewer comment:





1089  
 1090 *Figure 7: Imaginary part of the OA refractive index at 370 nm, obtained from offline UV/vis*  
 1091 *spectroscopy of methanol OA extracts, plotted as a function of  $f_{OA}$ . The data could be*  
 1092 *empirically represented by a linear function in the log-log space, in the measurement range.*  
 1093 *The ordinary least-squares fit is  $(k_{OA, nm}) = \log(M_{BC}/M_{OA})(0.51 \pm 0.07) + (-0.98 \pm 0.05)$ .*  
 1094 *Also shown are parameterizations of  $k_{OA}(370nm)$  for open burning against  $M_{BC}/M_{OA}$*   
 1095 *estimated based on the online  $k_{OA}(550nm)$  measurements in Saleh et al. (2014) and Lu*  
 1096 *et al. (2015), using the  $k_{OA}$  wavelength dependence reported by the respective authors.*

1097  
 1098 **Reference added:**

1099  
 1100 *Saleh, R., Robinson, E. S., Tkacik, D. S., Ahern, A. T., Liu, S., Aiken, A. C., Sullivan, R.*  
 1101 *C., Presto, A. A., Dubey, M. K., Yokelson, R. J., Donahue, N. M., and Robinson,*  
 1102 *A.L.: Brownness of organics in aerosols from biomass burning linked to their black*  
 1103 *carbon content, Nat. Geosci., 7, 2–5, doi:10.1038/ngeo2220, 2014.*  
 1104  
 1105  
 1106  
 1107  
 1108  
 1109  
 1110  
 1111  
 1112  
 1113  
 1114

1115

1116 27. Fig. S7: It is not clear why the propagated uncertainty in the AAE increases with  
1117 wavelength or  $f_{OA}$ . The AAE is a measured quantity that depends only on the measured  
1118 absorption at two wavelengths. . Why would uncertainty in absorption depend on  $f_{OA}$ ? And  
1119 are the authors saying that the uncertainty in absorption increases with wavelength?  
1120 Uncertainty in the AAE should directly propagate from Eqn. 10, which is independent of  
1121  $f_{OA}$ . Perhaps I am misunderstanding?

1122

1123 We believe that there is a misunderstanding. Fig. S7A is obtained from the error propagation  
1124 of equation 13 solved for different wavelengths, using the geometric mean and standard  
1125 deviation of  $MAC_{POA}(\lambda)$  and  $MAC_{BC}(\lambda)$ . The resulting error term represents the variability  
1126 in/ the confidence level on the  $\alpha(t_0, \lambda, 880nm)$  at different wavelengths. Equation 13 is  
1127 expressed below:

1128

1129 
$$\alpha(t_0, \lambda, 880nm) = \alpha_{BC+POA}(t_0, \lambda, 880nm)$$

$$= \frac{1}{\ln(880nm/\lambda)} \ln \left( \frac{MAC_{BC}(t_0, \lambda)}{MAC_{BC}(t_0, 880nm)} + \frac{M_{OA}(t_0)MAC_{POA}(t_0, \lambda)}{b_{abs}(t_0, 880nm)} \right)$$

1130

1131 As  $\alpha(t_0, \lambda, 880nm)$  depends on  $M_{OA}/b_{abs}(t_0, 880nm)$   $\sigma_{\alpha(t_0, \lambda, 880nm)}$  also does. We expressed  
1132  $M_{OA}/b_{abs}(880nm)$  as  $f_{OA}$ , using  $\sigma_{ATN}$  to estimate EC mass from  $b_{ATN}(880nm)$ .

1133 The image plot in panel B shows that at short wavelengths and low fractions of OA, the  
1134 confidence level on  $\alpha$  is within 0.1. However, with increasing  $f_{OA}$ , and at longer wavelength  
1135 the uncertainty in predicting  $\alpha$  increases. The idea behind this figure is to provide an error on  
1136 the predicted  $\alpha$  when the  $f_{OA}$  is extrapolated to values higher than measured here (typical of  
1137 open burning).

1138 We have updated the figure caption adding the explanations above:

1139 *Figure S7: Analysis of the fitting errors of  $\alpha(\lambda, 880nm)$  of primary emissions as a*  
1140 *function of  $f_{OA}$ . Panel A shows the  $\alpha$  residual as a probability density function. Panel*  
1141 *B is an image plot of the  $\alpha(\lambda, 880nm)$  error,  $\sigma_{\alpha(t_0, \lambda, 880nm)}$ , as a function of  $f_{OA}$  at*  
1142 *different wavelengths.  $\sigma_{\alpha(t_0, \lambda, 880nm)}$  is obtained from the error propagation of Eq.*  
1143 *(13) solved for different wavelengths, using the geometric mean and standard*  
1144 *deviation of  $MAC_{POA}(\lambda)$  and  $MAC_{BC}(\lambda)$ . This error term represents the variability in*  
1145 *or the confidence level on the  $\alpha(t_0, \lambda, 880nm)$  at different wavelengths. As*  
1146  *$\alpha(t_0, \lambda, 880nm)$  depends on  $M_{OA}/b_{abs}(t_0, 880nm)$  in Equation 13,  $\sigma_{\alpha(t_0, \lambda, 880nm)}$  also*  
1147 *does. We expressed  $M_{OA}/b_{abs}(t_0, 880nm)$  as  $f_{OA}$ , using  $\sigma_{ATN}$  to estimate EC mass*  
1148 *from  $b_{ATN}(880nm)$ . At short wavelengths and low OA fractions, the confidence level*  
1149 *on  $\alpha$  is within 0.1. However, with increasing  $f_{OA}$ , and at longer wavelength the*  
1150 *uncertainty in predicting  $\alpha$  increases.*

1151

1152

1153  
1154  
1155  
1156  
1157  
1158  
1159  
1160  
1161  
1162  
1163  
1164  
1165  
1166  
1167  
1168  
1169  
1170  
1171  
1172  
1173  
1174  
1175  
1176  
1177  
1178  
1179  
1180  
1181  
1182  
1183  
1184  
1185  
1186  
1187  
1188  
1189  
1190  
1191  
1192  
1193  
1194  
1195  
1196

28. L384: A larger GSD does not necessarily mean a larger uncertainty. This could be variability that is independent of uncertainty. I do not regard this as a true assessment of “uncertainty.” It is only an assessment of variability. The authors should, however, consider uncertainties explicitly.

The reviewer is correct, the word “variability” rather than “uncertainty” should have been used in this sentence. That is, the GSD values relate to variability in the MAC values that is not explained by the variability in  $f_{OA}$ . Overall, our data show that this variability is related to a real change in the chemical nature of the compounds present and their intrinsic absorptivity, as online MAC values correlate well with  $k_{OA}$  values independently measured offline after methanol extraction. As mentioned above, we have now added a new section discussing the quantifiable and unquantifiable uncertainties. We have also updated the related to the variability in the determined parameters as follows:

*Uncertainties and variability in  $MAC_{BC}$ ,  $MAC_{POA}$  and  $MAC_{SOA}$ . Table 1 shows the fitting errors related with  $MAC_{BC}(\lambda)$ ,  $MAC_{POA}(\lambda)$  and  $MAC_{SOA}(\lambda)$ , arising from our measurement precision and experimental variability. These fitting errors are greater than our estimated uncertainties in the absorption coefficients measured by MWAA (10%), and comparable to our estimated uncertainty in OA mass measured by AMS (30%). The residuals in the fitted  $MAC_{BC}(\lambda)$  are relatively low ( $< 10\%$ ), increasing with decreasing  $\lambda$ . By contrast, the uncertainties in the fitted  $MAC_{POA}(\lambda)$  are much higher ( $GSD = 1.2-1.5$ ) and increase with increasing  $\lambda$ . The relative residuals between the measured and fitted  $\alpha(\lambda, 880\text{nm})$  for primary emissions showed a mean bias and RMSE of 0.07 and 0.13, respectively (Fig. S7), indicating that our fitted MAC results provide a good description of the data set.  $MAC_{SOA}(\lambda)$  values determined were highly variable between experiments with a  $GSD = 1.39$  and  $2.42$  for  $\lambda=370$  nm and  $660$  nm, respectively. In Fig. S10, we show the distribution of  $MAC_{SOA}(\lambda)$  values as box and whiskers against OH exposure, showing no particular dependence of these values with aging as it will be discussed below. Therefore, we expect the fitting errors in  $MAC_{SOA}$  and of  $MAC_{POA}$  to be mainly related to true changes in the organic aerosol chemical composition between different burns, since the variability of  $MAC_{BC}(\lambda)$  was relatively small. In Section 4.3, we discuss this variability further using the results of an additional and independent analysis.*

1197  
1198  
1199  
1200  
1201  
1202  
1203  
1204  
1205  
1206  
1207  
1208  
1209  
1210  
1211  
1212  
1213  
1214  
1215  
1216  
1217  
1218  
1219  
1220  
1221  
1222  
1223  
1224  
1225  
1226  
1227  
1228  
1229  
1230  
1231  
1232  
1233  
1234  
1235  
1236  
1237  
1238  
1239  
1240

29. L386: Given that the authors show distributions and fits for the AAE, it would seem appropriate to also show similar for the MAC\_POA and MAC\_SOA so that the reviewer can judge. Given the width of the SOA distribution, is a normal fit even appropriate? (Probably not, in a fundamental sense, since MAC values cannot be  $< 0$ . But perhaps a normal distribution is appropriate in a practical sense.)

The MACs of primary species (POA and BC) at different wavelengths are obtained through fitting equation 13, which used the AAE as dependent variable. Therefore, we have assessed the model goodness of fit by showing the residuals in the AAE values (Fig. S7). While we cannot show the residuals distributions for  $MAC_{POA}$  and  $MAC_{BC}$  resulting from the fit, the obtained fitting errors (GSD) can be used to represent these distributions. The normal fit of the AAE residuals in Fig. S7A serves only to illustrate the distribution of fitting errors and is not essential to our analysis. We note that we have never claimed that MAC values to be normally distributed and hence we performed our fits in log-transformed space to constrain the MAC to be greater than zero.  $MAC_{SOA}$  is the only unknown parameter in equation 19. Therefore, we have shown the distributions of  $MAC_{SOA}$  as a function of wavelength and OH exposure, in Fig. 4 and 5, respectively.

30. L398: The authors should provide the resulting uncertainties, or at least ranges, based on the multivariate analysis, for the AAE values.

The ranges for AAE values have been already provided in Fig. S9 and Table S1, and following the reviewer's comment we have added the following to the main text:

*“This yielded  $\alpha_{BC} = 1.2$ , [...], with corresponding uncertainties of approximately 20% (complete details of the uncertainties are provided in Table S1).”*

31. L402: I suggest removing the “this is the first study” statement. Saleh (2014) reported very closely related “w” values, from which AAE values can be calculated, for SOA from biomass burning.

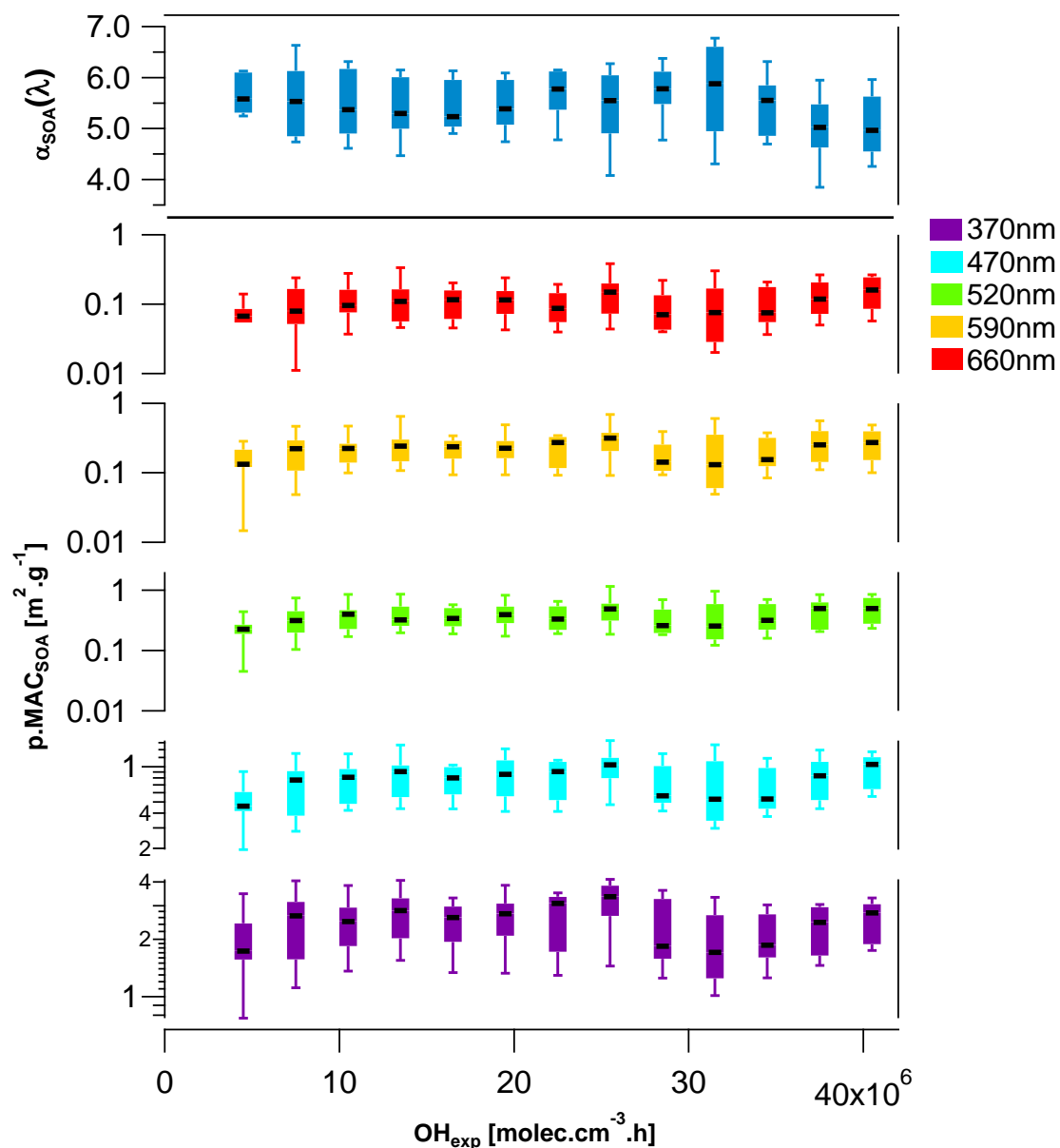
The new sentence reads as follows:

*The high  $\alpha$  values obtained for the organic fractions are consistent with previous measurements for BrC containing POA (e.g. Chakrabarty et al., 2010, 2013).*

1241  
1242  
1243  
1244  
1245  
1246  
1247  
1248  
1249

32. Fig. S10: I find that the use of the log scale for the y-axis makes it difficult for the reader to see what sort of changes did/did not occur. Variability in the AAE over so many orders of magnitude is not expected, but a factor of 2 would be reasonable. Thus, a linear scale should be used.

In response to the reviewer comment, we have modified the figure to better illustrate the variability in our data.



1250  
1251  
1252  
1253  
1254  
1255  
1256

1257  
1258  
1259  
1260  
1261  
1262  
1263  
1264  
1265  
1266  
1267  
1268  
1269  
1270  
1271  
1272  
1273  
1274  
1275  
1276  
1277  
1278  
1279

33. L424: In Fig. S13, and Fig. 6, it is unclear why the authors fit only the “aged” data. Why exclude the primary, especially in Fig. S13? Because the relationship is visibly much worse? This goes to the statement about sensitivity to Mie calculations.

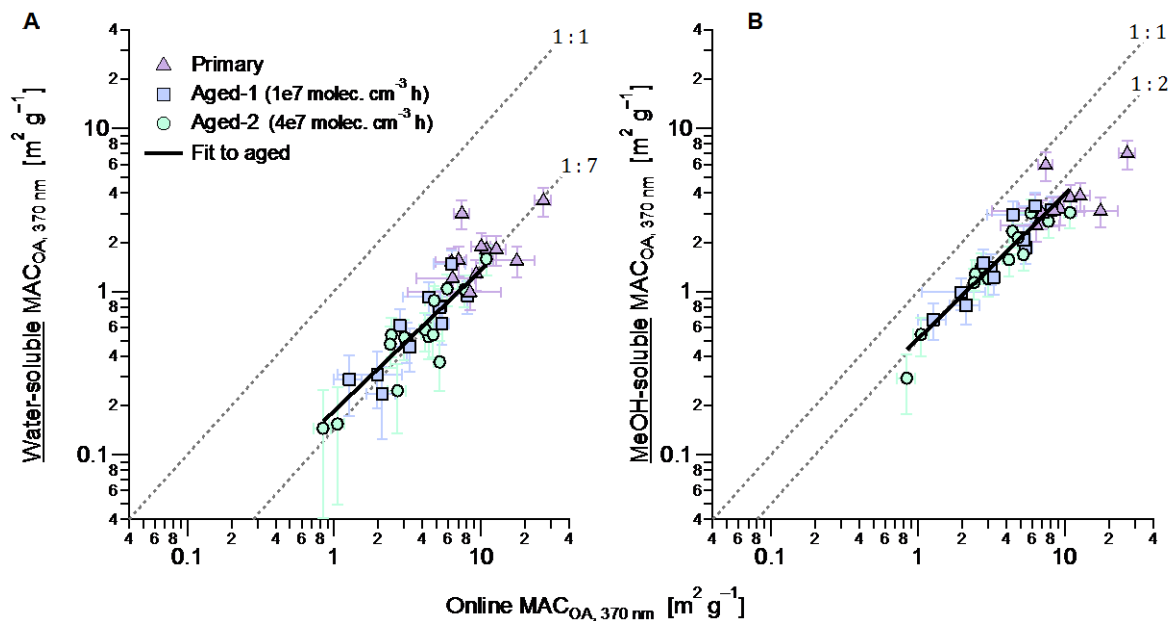
We apologize; this information went missing during revisions of an earlier manuscript draft. We excluded the primary because our assumption of particle sphericity inherent in the Mie calculations is generally more likely to be violated for fresh than for aged combustion particles. Similarly, our interpretation of the SMPS-measured mobility diameter as representative of a physical particle diameter is violated in the presence of fractal-like soot particles, which have a shape factor significantly different from unity.

The variability in our primary results is interpreted as illustrating directly the impact of this variability on our analysis. This variability will reflect differences in burn conditions as well as the chaotic impacts of the combustion process (for example, uncontrolled differences between time spent in pre-ignition [where most OA is emitted] versus flaming phases [where most BC is emitted], between the physical distribution of flames during each burn, etc.).

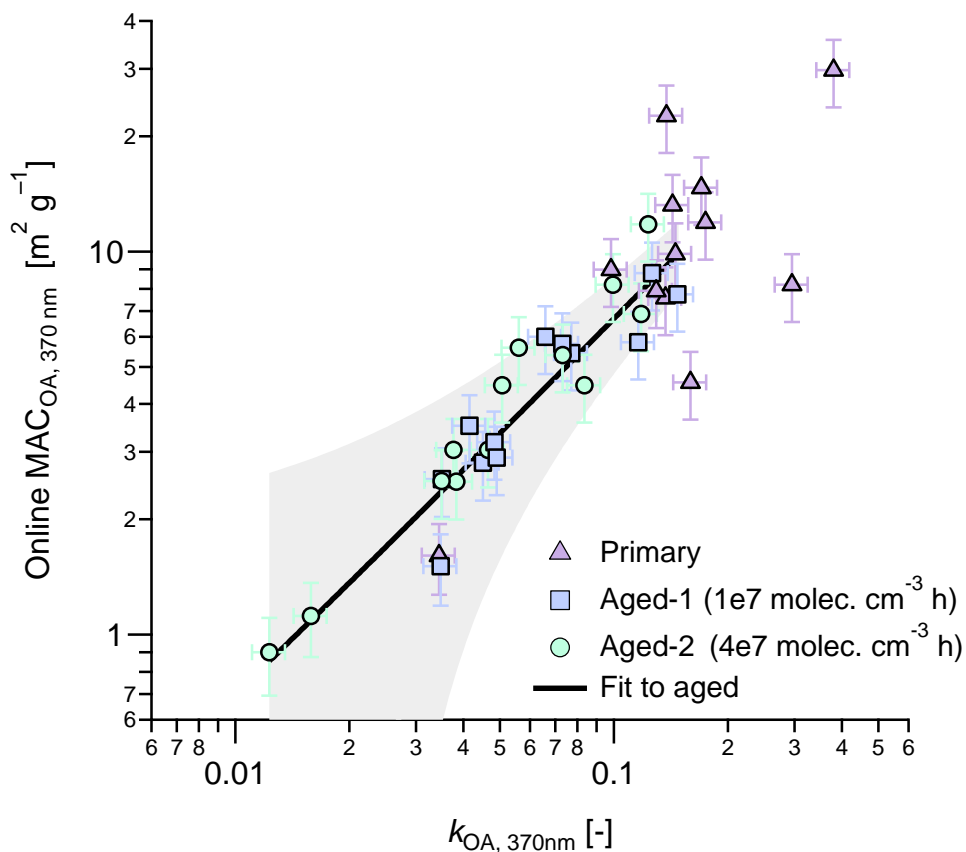
In spite of these simplifications, the fresh data in Fig. 6 generally fall on the best-fit line for the aged data. The fact that they fall on the best-fit line is a direct demonstration of the magnitude of the uncertainties in our retrieved OA MAC.

We have revised the figures and captions for clarity. The revised figures are:

1280



1281



1282

1283

1284 We modified the text:

1285

1286

1287

1288

1289

L431. The data in Fig. 6B show that the methanol extracts correspond to a MAC about 50% smaller than the online data. The scatter in the data is significantly reduced for the aged data (note that, in this analysis, aged OA refers to the sum of POA and SOA, since the reported values represent all OA after aging). This reduced

1290 scatter is expected, considering that aging is likely to result in more-spherical  
1291 particles. We have assumed particle sphericity when interpreting the SMPS data and  
1292 performing the Mie analysis. While the propagation of quantifiable uncertainties  
1293 leads to an error estimate of ~25%, considering the simplifications that were necessary  
1294 for the Mie analysis, we consider a 50% closure to be an adequate agreement.  
1295 Despite this, we cannot exclude additional methanol insoluble brown carbon.  
1296 Conversely, the fit in Figure 6A indicates that the apparent MAC of water-soluble  
1297 species was a fourth of the respective methanol MAC, according to the slope of only  
1298  $12 \pm 3\%$ . Only the aged data have been fit to illustrate this point. This strong  
1299 disagreement shows that the BrC in our samples was hardly water soluble, even for  
1300 the most aged samples. As we expect that the majority of OA in our samples formed by  
1301 wood pyrolysis (Di Blasi, 2008; Corbin et al., 2015b; Shafizadeh, 1984), we can  
1302 compare our results directly to those of Chen and Bond (2010), who also found that  
1303 primary wood-pyrolysis BrC was water insoluble. Moreover, the water-insoluble  
1304 nature of the light absorbing components of SOA is in line with the results by Bruns et  
1305 al. (2016) who showed that SOA precursors during these experiments were  
1306 predominantly aromatic compounds.

1307

1308 We modified the Figure 6 caption:

1309

1310 *Figure 6: Comparison of the  $MAC_{OA}(370nm)$  of aged aerosols determined from online  
1311 and offline absorption measurements. The offline filter-extraction method directly  
1312 quantified properties of total OA (ordinate), while the average of  $MAC_{SOA}$  and  
1313  $MAC_{POA}$  weighted with respective mass concentrations is shown on the abscissa. The  
1314 panels show offline measurements of (A) water-soluble OA, (B) methanol-soluble OA.  
1315 Fits are to aged data only due to the significantly smaller scatter of those data,  
1316 although primary data on average follow similar trends. The fitted slopes and  
1317 intercepts are, respectively, (A)  $0.13 \pm 0.02$  and  $0.05 \pm 0.06 \text{ m}^2\text{g}^{-1}$  and (B)  $0.12 \pm 0.1$   
1318 and  $0.38 \pm 0.03 \text{ m}^2\text{g}^{-1}$ .*

1319

1320 We modified the Figure S13 caption:

1321

1322 *Figure S13:  $MAC_{OA}$  at  $\lambda = 370 \text{ nm}$  calculated from aethalometer measurements vs.  
1323  $k_{OA}$  at  $\lambda = 370 \text{ nm}$  from the UV/visible measurements of the methanol extracts. The  
1324 shaded region shows the 90% confidence interval of a weighted orthogonal  
1325 regression (slope  $66 \pm 9 \text{ m}^2\text{g}^{-1}$ , intercept  $0.0 \pm 0.3 \text{ m}^2\text{g}^{-1}$ ) to illustrate the relatively  
1326 small range of variability in the data for aged samples.*

1327

1328

1329

1330

1331

1332

1333

1334



1335  
1336  
1337  
1338  
1339  
1340  
1341  
1342  
1343  
1344  
1345  
1346  
1347  
1348  
1349  
1350  
1351  
1352  
1353  
1354  
1355  
1356  
1357  
1358  
1359  
1360  
1361  
1362  
1363  
1364  
1365  
1366  
1367  
1368  
1369  
1370  
1371  
1372  
1373  
1374  
1375  
1376  
1377

34. L427: Fig. 6 normalizes out any uncertainty/variability in the measured [OA], because both absorption values are normalized by this. Fig. 4, in contrast, does not. How can the authors rule out the possibility that there is not some variability in the measurement of OA between burns, perhaps dependent on particle shape or variability in bounce in the AMS (which can differ between POA and SOA)?

Fig. 4 actually does normalize the absorption by [OA], so we are not sure which figure the reviewer had in mind. Our goal with Fig. 6 is to relate the offline and online absorption measurements, in such a way that any unknown uncertainties would influence scatter in the plot.

The bounce-related collection efficiency of the AMS was concluded as close to 1.0 for wood-burning OA in the literature reviewed by Corbin et al. (2015b; in their Section S1.2). From recent results from our group using a similar setup, we have measured with an AMS primary organic aerosol rich emissions from smoldering biomass and found the POA collection efficiency to be close to 1.

We reanalyzed our SMPS, AMS, and eBC (MWAA-calibrated AE33) data for the present study by fitting the SMPS mass (predicted with a density of 1.5 g/cm<sup>3</sup>) against the total PM mass predicted as AMS OA + eBC. The 95% CI of the slope of this fit corresponds to a CE of 0.7-1.0 (relative 19% relative uncertainty), consistent with the literature cited above. Combining this 19% uncertainty (slope uncertainty) with the 30% uncertainty already assigned to the AMS OA (largely reflecting uncertainties in RIE) in quadrature results in a 36% uncertainty in AMS OA, which we have updated in the text.

Shape-related collection efficiency issues in the AMS are unlikely as such issues mainly come into play when transmission through the AMS lens is considered. That is, shape itself is not an issue in the AMS except as it affects aerodynamic diameters (DeCarlo et al., 2014). In our study, particles were large enough that transmission and therefore shape-related issues were minor.

We modified the text:

*L94. Uncertainties related to particle collection efficiency in the AMS are considered negligible for the relatively-large particles sampled here, which in terms of volume are within the size range transmitted efficiently by the AMS aerodynamic lens (Liu et al., 2007). The collection efficiency of wood-combustion OA is expected to be unity (Corbin et al., 2015b).*

1378  
1379  
1380  
1381  
1382  
1383  
  
1384  
1385  
1386  
1387  
1388  
  
1389  
1390  
1391  
1392  
1393  
1394  
1395  
1396  
1397  
1398  
1399  
1400  
1401  
1402  
1403  
1404  
1405  
1406  
1407  
1408  
1409  
1410  
1411  
1412  
1413  
1414  
1415  
1416  
1417  
1418  
1419  
1420  
1421

**References:**

*Corbin, J. C., Keller, A., Lohmann, U., Burtscher, H., Sierau, B. and Mensah, A. A.: Organic emissions from a wood stove and a pellet stove before and after simulated atmospheric aging, Aerosol Sci. Technol., 49(11), 1037–1050, doi:10.1080/02786826.2015.1079586, 2015b.*

*DeCarlo, P. F., Slowik, J. G., Worsnop, D. R., Davidovits, P., and Jimenez, J. L.: Particle morphology and density characterization by combined mobility and aerodynamic diameter measurements. Part 1: Theory, Aerosol Sci. Technol., 38, 1185–1205, 2004.*

*Liu, P. S. K., Deng, R., Smith, K. A., Jayne, J. T., Williams, L.R., Canagaratna, M. R., Moore, K., Onasch, T. B., Worsnop, D.R., and Deshler, T.: Transmission efficiency of an aerodynamic focusing lens system: comparison of model calculations and laboratory measurements for the aerodyne aerosol mass spectrometer, Aerosol Sci. Tech., 41, 721–733, 2007.*

35. L432: This 46% must state that it is for aged OA only. It remains unclear to me why the primary is excluded.

This statement was removed following comment #33.

36. L440: Is this a fair comparison, given that the authors have focused on the aged OA?

This comment is deprecated following our response to comment #33, but we would also like to point out that we did point out in the text that Chen and Bond studied primary and not aged OA.

37. L431: Are these fits forced through zero?

No fits in this work were forced through zero, but all intercepts were not significantly different from zero. We acknowledge that it is our responsibility to report those intercepts clearly and have updated the text where fits are still reported (the present fits are no longer discussed, see response to comment 33).

38. L441: The authors seem to be implying that SOA formed from oxidation of aromatic precursors is not especially water soluble, or at least less soluble than in methanol. The authors might consider citing e.g. (*Zhang et al., 2011*), to strengthen this argument.

Thank you, we have added this citation in the corrected version of the manuscript.

1422  
1423  
1424  
1425  
1426  
1427  
1428  
1429  
1430  
1431  
1432  
1433  
1434  
1435  
1436  
1437  
1438  
1439  
1440  
1441  
1442  
1443  
1444  
1445  
1446  
1447  
1448  
1449  
1450  
1451  
1452  
1453  
1454  
1455  
1456  
1457  
1458  
1459  
1460  
1461  
1462  
1463  
1464

**Reference:**

Zhang, X., Lin, Y. -H., Surratt, J. D., Zotter, P. and Prévôt, A. S. H.: Light-absorbing soluble organic aerosol in Los-Angeles and Atlanta: A contrast in secondary organic aerosol, *Geophys. Res. Lett*, 38, 2011.

39. Fig. 7: As already noted above, I find the comparison here insufficient. Saleh et al. (2014) and Lu et al. (2015) report values not at 370 nm. This is unclear. Also, the line shown for Lu et al. (2015) appears to be incorrect. See their Fig. 1D. Further, and importantly, the Lu et al. (2015) data are largely, although not entirely, derived from the Saleh measurements. Thus, they are not really an independent assessment.

We have replied to comment #26 about the data in Fig. 7 and we think we have addressed all the points raised by the reviewer in this question. We acknowledge that the Saleh and Lu data sets are not entirely independent, but have treated the corresponding parameterizations as independent for lack of any method to disentangle their interdependence. We expect other studies to cite Saleh and Lu's separately. Note that we have not focused on the more comprehensive Lu data set because the Saleh dataset represents biomass burning specifically, which we have also studied in the present work.

*Figure 7: Imaginary part of the OA refractive index at 370 nm, obtained from offline UV/vis spectroscopy of methanol OA extracts, plotted as a function of fOA. The data could be empirically represented by a linear function in the log-log space, in the measurement range. The ordinary least-squares fit is  $(k_{OA, nm}) = \log(M_{BC}/M_{OA})(0.51 \pm 0.07) + (-0.98 \pm 0.05)$ . Also shown are parameterizations of  $k_{OA}(370nm)$  for open burning against  $M_{BC}/M_{OA}$  estimated based on the  $k_{OA}(550nm)$  measurements in Saleh et al. (2014) and Lu et al. (2015), using the  $k_{OA}$  wavelength dependence reported by the respective authors.*

We have also updated the text at line 460:

*The parameterizations reported by these authors are included in Fig. 7, where the wavelength dependence reported by those authors has been used to adjust their parameterizations to 370nm.*

1465  
1466  
1467  
1468  
1469  
1470  
1471  
1472  
1473  
1474  
1475  
1476  
1477  
1478  
1479  
  
1480  
1481  
1482  
1483  
1484  
1485  
1486  
1487  
1488  
1489  
  
1490  
1491  
1492  
1493  
1494  
1495  
1496  
1497  
1498  
1499  
1500  
1501  
1502  
1503  
1504  
1505  
1506

40. Fig. 7: The logic of a linear fit to the observations is not clear to me. The authors have argued that the SOA is absorbing, and differently absorbing than the POA. If I use the equation given and extrapolate towards  $M_{BC}/M_{OA} \rightarrow 0$ , the  $k_{OA} \rightarrow 0$ . If the SOA is absorbing, and if SOA formation drives the decrease in the  $M_{BC}/M_{OA}$ , then the limiting value of  $k_{OA}$  should be equal to the value for  $k_{SOA}$ . As such, the provided fit does not seem appropriate and requires justification. Some of this may be experiment-to-experiment variability. But the limiting case issue remains.

We do agree with the reviewer that  $k_{OA}$  will likely tend towards  $k_{POA}$  and  $k_{SOA}$  when  $M_{BC}/M_{OA}$  is very large and very small, respectively. However, within the range covered, a line is the simplest model which can describe our data adequately and the linear fit used is simply empirical. We do not aim to include a complete physical interpretation in this simple fit. With our fit in Fig. 7, we mainly wish to emphasize that  $k_{OA}$  can be described as a function of  $M_{BC}/M_{OA}$  only. For clarification we have modified the figure caption as follows:

*Figure 7: Imaginary part of the OA refractive index at 370 nm, obtained from offline UV/vis spectroscopy of methanol OA extracts, plotted as a function of  $f_{OA}$ . The ordinary least-squares fit is  $\log(k_{OA,370nm}) = \log(M_{BC}/M_{OA})(0.51 \pm 0.07) + (-0.98 \pm 0.05)$  and illustrates that the observed  $k_{OA}$  can be described as a function of  $M_{BC}/M_{OA}$  with reasonable accuracy, regardless of the degree of aging. Also shown are parameterizations of  $k_{OA}(370nm)$  for open burning against  $M_{BC}/M_{OA}$  estimated based on the online  $k_{OA}(550nm)$  measurements in Saleh et al. (2014) and Lu et al. (2015), using the  $k_{OA}$  wavelength dependence reported by the respective authors.*

41. Fig. 7: The authors should be able to, from their observations and within their assumptions, calculate  $M_{POA}/M_{SOA}$ . They might consider plotting  $k_{OA}$  vs. this ratio instead of versus  $M_{BC}$ . These will be related, of course, since the authors assume POA is proportional to BC during aging for a given experiment.

We agree, however, we do not think that this calculation will shed any new insights into our data set. In Fig. 7, we have only chosen to use  $M_{BC}/M_{OA}$  as an abscissa because previous studies have used this quantity; our goal in Fig. 7 is a comparison of our results with related literature. We do not find this ratio to be particularly meaningful or interesting physically, but we acknowledge that future studies are likely to measure it as well and so it provides a useful basis for comparison.

1507  
1508  
1509  
1510  
1511  
1512  
1513  
1514  
1515  
1516  
1517  
1518  
1519  
1520  
1521  
1522  
1523  
1524  
1525  
1526  
1527  
1528  
1529  
1530  
1531  
1532  
1533  
1534  
1535  
1536  
1537  
1538  
1539  
1540  
1541  
1542  
1543  
1544  
1545  
1546  
1547  
1548

42. The origin of these “uncertainties” is unclear. They are explained later for  $f_{OA}$ , but for the MAC values it is not abundantly clear.

We apologize but we do not see which part of the manuscript the reviewer is referring to. We assume that the reviewer is referring to the fitted MAC values, which we have commented on above and adjusted the manuscript to include.

43. L486: This statement regarding mass yields of SOA requires much further detail.

The analysis of SOA gas phase precursors has been thoroughly presented in Bruns et al. (2016) and the discussion about SOA yields is beyond the scope of this study. As this statement is not required for the understanding of the paper we have removed it in the corrected version of the manuscript. The section now reads as follows:

*The  $M_{SOAP,WLC}/M_{POA,WLC}$  was on average equal to 7.8 (GSD = 1.4) and  $k_{OH}$  was estimated as  $2.7 \times 10^{-11}$  molecule<sup>-1</sup> cm<sup>3</sup> (GSD = 1.4), consistent with the SOA precursors chemical nature measured (e.g. PAH and phenol derivatives) by a proton-transfer-reaction mass spectrometer (PTR-MS) (Bruns et al., 2016, 2017). These high rates and enhancement ratios indicate the rapid production of SOA.*

44. L512: A note about terminology. I am not certain that “error analysis” is appropriate here. Variance in the POA fraction is not “error.” It is variability. A substantial aspect of this “error analysis” is really just a “sensitivity analysis.” I suggest that the authors limit the term “error analysis” to when they are truly considering errors, and use some other term when they are considering variability. This is true here and elsewhere.

We agree with the reviewer and have changed the word “error” to “sensitivity”.

We have also made the requested modifications related with the section entitled “*Uncertainties and variability in  $MAC_{BC}$ ,  $MAC_{POA}$  and  $MAC_{SOA}$* ”.

45. L499: The authors should clarify the origin of the solar irradiance data that they have used.

We have now added the reference on which the solar irradiance data are based:

*Gueymard, C.; Myers, D.; Emery, K. "Proposed Reference Irradiance Spectra for Solar Energy Systems Testing," Solar Energy, 73, 6, 443–467, 2002.*

# 1 **Production of particulate brown carbon during atmospheric** 2 **aging of residential wood-burning emissions**

3 Nivedita K. Kumar<sup>1</sup>, Joel C. Corbin<sup>1\*</sup>, Emily A. Bruns<sup>1</sup>, Dario Massabó<sup>2</sup>, Jay G. Slowik<sup>1</sup>, Luka  
4 Drinovec<sup>3,4</sup>, Griša Močnik<sup>3,4</sup>, Paolo Prati<sup>2</sup>, Athanasia Vlachou<sup>1</sup>, Urs Baltensperger<sup>1</sup>, Martin  
5 Gysel<sup>1</sup>, Imad El-Haddad<sup>1</sup> and André S. H. Prévôt<sup>1</sup>

6 <sup>1</sup>Laboratory of Atmospheric Chemistry, Paul Scherrer Institute, 5232 Villigen, Switzerland

7 <sup>2</sup>Department of Physics & INFN, University of Genoa, via Dodecaneso 33, 16146, Genova, Italy

8 <sup>3</sup>Aerosol d.o.o, Kamniška 41, 1000 Ljubljana, Slovenia

9 <sup>4</sup>Condensed Matter Physics, Jožef Stefan Institute, 1000 Ljubljana, Slovenia

10 \*Now at National Research Council Canada, Ottawa, Canada

11 *Correspondence to:* I. El-Haddad ([imad.el-haddad@psi.ch](mailto:imad.el-haddad@psi.ch)), A. S. H. Prévôt ([andre.prevot@psi.ch](mailto:andre.prevot@psi.ch))

## 12 **ABSTRACT**

13 We investigate the optical properties of light-absorbing organic carbon (brown carbon) from domestic wood  
14 combustion as a function of simulated atmospheric aging. ~~At shorter wavelengths, light absorption by brown carbon~~  
15 ~~from primary organic aerosol (POA) and secondary organic aerosol (SOA) formed during aging was around 10 %~~  
16 ~~and 20 %, respectively, of the total aerosol absorption (brown carbon plus black carbon).~~ **At shorter wavelengths**  
17 **(370 – 470nm), light absorption by brown carbon from primary organic aerosol (POA) and secondary organic**  
18 **aerosol (SOA) formed during aging was around 10 % and 20 %, respectively, of the total aerosol absorption (brown**  
19 **carbon plus black carbon).** The mass absorption cross-section (MAC) determined for black carbon (BC, 13.7 m<sup>2</sup> g<sup>-1</sup>  
20 (geometric standard deviation GSD = 1.1) at 370 nm) was consistent with that recommended by Bond et al. (2006).  
21 The corresponding MAC of POA (5.5 m<sup>2</sup> g<sup>-1</sup> (GSD = 1.2)) was higher than that of SOA (2.4 m<sup>2</sup> g<sup>-1</sup> (GSD = 1.3)) at

22 370 nm. However, SOA presents a substantial mass fraction, with a measured average SOA/POA mass ratio after  
23 aging of ~5 and therefore contributes significantly to the overall light absorption, highlighting the importance of  
24 wood-combustion SOA as a source of atmospheric brown carbon. The wavelength dependence of POA and SOA  
25 light absorption between 370 nm and 660 nm is well described with absorption Ångström exponents of 4.6 and 5.6,  
26 respectively. UV-visible absorbance measurements of water and methanol-extracted OA were also performed  
27 showing that the majority of the light-absorbing OA is water insoluble even after aging.

28

## 29 1. INTRODUCTION

30 Atmospheric aerosols contribute to radiative forcing either directly by absorbing and scattering light or indirectly by  
31 acting as cloud-condensation and ice nuclei. While black carbon (BC) from combustion processes is the most  
32 efficient light-absorbing aerosol component, organic aerosols (OA) may also absorb solar radiation (Alexander et  
33 al., 2008; Chen and Bond, 2009; Kirchstetter et al., 2004). This light-absorbing OA, denoted as brown carbon  
34 (BrC), absorbs most strongly at shorter UV-visible wavelengths (Andreae and Gelencsér, 2006; Hoffer et al., 2005).  
35 Global chemical-transport model estimates indicate that the BrC contribution to the positive radiative forcing of  
36 climate by anthropogenic aerosols may not be negligible (Feng et al., 2013; Jo et al., 2016; Lin et al., 2014; Wang et  
37 al., 2014).

38 Unlike BC, whose light absorption properties are relatively constant across sources (Bond et al., 2013), BrC is  
39 composed of a wide range of largely unknown compounds, which exhibit highly variable spectral dependence and  
40 absorption efficiencies. For example, reported imaginary indices of refraction for different organic species, which  
41 describe the absorption of these compounds, span two orders of magnitude (Lu et al., 2015). Because it is  
42 impractical to experimentally separate BrC from non-absorbing OA, optical properties are typically determined for  
43 the bulk OA of a given source. The large variability of BrC fraction in combustion aerosol may contribute to the  
44 wide variation in reported properties of BrC containing OA.

45 Biomass burning OA, which contributes two-thirds of the global budget of directly-emitted primary OA (POA), is  
46 expected to be a considerable source of BrC (Chakrabarty et al., 2010; Hecobian et al., 2010; Lack and Langridge,

47 2013; Liu et al., 2014). The variability in reported light absorption properties of biomass burning OA with fuel type  
48 and burn conditions remains a major obstacle complicating its treatment in climate models (Lu et al., 2015; Saleh et  
49 al., 2013). Residential biomass burning is typically characterized by a more efficient combustion, than open burning.  
50 Residential wood burning represents a substantial contribution to anthropogenic combustion emissions (Bond et al.,  
51 2013), especially in urban atmospheres, and is considered the largest source of OA in Europe during winter (Denier  
52 Van Der Gon et al., 2015).

53 Upon photo-oxidation, biomass-burning emissions produce secondary organic aerosol (SOA) at concentrations  
54 similar to or exceeding the primary organic aerosol (POA) (Bertrand et al., 2017; Bruns et al., 2015, 2016; Corbin et  
55 al., 2015a; Grieshop et al., 2009). There is a growing body of evidence that light absorption by OA change with OH  
56 exposure (aging) owing to the production of secondary BrC or to the transformation of primary BrC (Forrister et al.,  
57 2015; Heringa et al., 2011; Lee et al., 2014; Zhao et al., 2015). However, these effects have not yet been  
58 systematically investigated and must be quantified to assess the climate effects of primary and aged biomass burning  
59 OA.

60 Here, we show that both POA and SOA from residential biomass burning emissions aged in controlled smog  
61 chamber experiments contain BrC. Wavelength dependent, mass-normalized absorption cross-sections (MACs) of  
62 POA and SOA are presented from online aerosol measurements as a function of aging for the first time.  
63 Complementary measurements of filter-extract absorbance (conducted in different solvents) are used to obtain the  
64 imaginary refractive index and to investigate the solubility of BrC in fresh and aged OA. While results presented  
65 here are related to flaming residential wood combustion emissions and cannot therefore be generalized, the approach  
66 used can be extrapolated for the characterization and quantification of the contribution of BrC in other primary and  
67 aged emissions.

68

## 69 **2. METHODS**

### 70 **2.1 Smog chamber experiments**

71 ~~Laboratory measurements were conducted in an 8 m<sup>3</sup> Teflon smog chamber (Bruns et al., 2015; Platt et al., 2013)~~  
72 ~~installed within a temperature controlled housing. Conditions in the chamber were maintained to represent winter~~



73 time in Europe, i.e. relative humidity ranging between 50–90%, at 263 K (Bruns et al., 2015, 2016). Beech wood  
74 was combusted in a residential wood stove. Primary emissions were sampled through heated lines at 413 K, diluted  
75 by a factor of ~14 using an ejector diluter (DI-1000, Dekati Ltd.), then sampled into the chamber, which provided an  
76 additional ten-fold dilution. The overall dilution was a factor of 100 to 200. As we aimed to sample only flaming-  
77 phase emissions into the chamber, samples were taken when the modified combustion efficiency (ratio of CO<sub>2</sub> to the  
78 sum of CO and CO<sub>2</sub>) was > 0.90. Despite maintaining the same combustion conditions, the resulting organic fraction  
79 in the different samples was highly variable, indicating that these samples are representative of a mixture of pre-  
80 ignition and flaming emissions (with varying contributions of each combustion stage).

81 Laboratory measurements were conducted in an 8 m<sup>3</sup> Teflon smog chamber (Bruns et al., 2015; Platt et al., 2013)  
82 installed within a temperature-controlled housing. Conditions in the chamber were maintained to represent winter  
83 time in Europe, i.e. relative humidity ranging between 50 – 90%, at 263 K (Bruns et al., 2015, 2016). Beech wood  
84 was combusted in a residential wood stove. Primary emissions were sampled through heated lines at 413 K, diluted  
85 by a factor of ~14 using an ejector diluter (DI-1000, Dekati Ltd.), then sampled into the chamber, which provided an  
86 additional ten-fold dilution. The overall dilution was a factor of 100 to 200. As we aimed to sample only flaming-  
87 phase emissions into the chamber, samples were taken when the modified combustion efficiency (ratio of CO<sub>2</sub> to the  
88 sum of CO and CO<sub>2</sub>) was > 0.90. Despite maintaining the same combustion conditions, the resulting organic fraction  
89 to the total carbonaceous aerosols in the different samples was highly variable, indicating that these samples are  
90 representative of a mixture of pre- ignition and flaming emissions (with varying contributions of each combustion  
91 stage). Finally, the resulting NO<sub>x</sub>/NMOG ratios, which dramatically influence SOA formation through influencing  
92 the fate of peroxy radicals, RO<sub>2</sub>, were estimated to be between 0.035 – 0.35 ppm ppm C<sup>-1</sup> (Bruns et al., 2016). These  
93 conditions can be considered as high NO<sub>x</sub> representative of urban/sub-urban conditions, where most of the RO<sub>2</sub>  
94 radicals react with NO, rather with RO<sub>2</sub>/HO<sub>2</sub>.

95 After injection of the primary emissions and stabilization of the concentrations, nitrous acid (HONO) was  
96 continuously added, which dissociates upon irradiation ( $\lambda < 400$  nm) and forms the hydroxyl radical (OH). Then, 9-  
97 times deuterated butanol sample (butanol- D9, 98%, Cambridge Isotope Laboratories) was subsequently injected  
98 into the chamber. The decay of butanol-D9 was used to infer the time-resolved OH exposure of the sampled aerosol  
99 (Barnet et al., 2012). The chamber was exposed to UV lights for ~3.5 hours.

100 Particles were collected onto filters (47 mm Tissue-quartz, Pall Corporation, 26 L min<sup>-1</sup> for 30-32 min) for offline  
101 optical measurements and the determination of elemental carbon (EC) mass. Three filters were collected during each  
102 experiment, namely i) a primary aerosol filter sample (“primary”), ii) a slightly aged aerosol (“Aged1”, OH  
103 exposure ~ 1x10<sup>7</sup> molecules cm<sup>-3</sup> h), collected 30 minutes after the UV lights were switched on, and iii) an aged  
104 aerosol (“Aged2”, OH exposure ~ 4x10<sup>7</sup> molecules cm<sup>-3</sup> h), collected at the end of the experiment (see Fig. S1 for  
105 the sampling periods). A charcoal denuder was installed upstream of the filter sampler to remove organic gases.  
106 Filters were stored at 253K until analysis.

107 In addition to the characterization of the particle optical properties detailed in the next section, a set of online and  
108 offline techniques were used for the characterization of the gaseous and particulate emissions before and after aging.  
109 The non-refractory particle size-segregated chemical composition was measured with a high resolution (HR) time-  
110 of-flight aerosol mass spectrometer (AMS) (DeCarlo et al., 2006). **Uncertainties related to particle collection**  
111 **efficiency in the AMS are considered negligible for the relatively-large particles sampled here, which in terms of**  
112 **volume are within the size range transmitted efficiently by the AMS aerodynamic lens (Liu et al., 2007). The**  
113 **collection efficiency of wood-combustion OA is expected to be unity (Corbin et al., 2015b).** Details related to the  
114 AMS data analysis and calibration can be found elsewhere (Bruns et al., 2015, 2016). A scanning mobility particle  
115 sizer was used to measure the size distribution of the evolving aerosol. Organic gases were monitored by a proton  
116 transfer reaction time-of-flight mass spectrometer (PTR-MS, [H<sub>3</sub>O<sup>+</sup>] reagent ion, Ionicon Analytik GmbH) (Bruns et  
117 al., 2017), following the same procedure as in Klein et al. (2016). Additionally, elemental carbon (EC) mass  
118 concentration was measured offline using a sunset thermo-optical analyzer, following the EUSAAR2 protocol  
119 (Cavalli et al., 2010).

120

121

122

123

124

125

126

## 127 2.2 Optical measurements

128 **Aethalometer.** A dual-spot aethalometer (Magee Scientific aethalometer AE33, Aerosol d.o.o.) was used for real-  
129 time aerosol light attenuation measurements at seven wavelengths ( $\lambda = 370, 470, 520, 590, 660, 880$  and  $950$  nm)  
130 (Drinovec et al., 2015). The instrument measures the attenuation coefficient ( $b_{\text{ATN}}$ ) of a light beam transmitted  
131 through a filter tape loaded with aerosol samples. The use of the sampling flow (here,  $2 \text{ L min}^{-1}$ ), integration time for  
132 the measurement (here, 1 minute), and automated dual-spot loading compensation to obtain  $b_{\text{ATN}}$  has been described  
133 by Drinovec et al. (2015).

134 The loading compensated  $b_{\text{ATN}}$  was used to infer the aerosol absorption coefficient,  $b_{\text{abs}}$ , using a constant wavelength  
135 independent correction factor  $C$ , which accounts for multiple scattering within the filter matrix (Weingartner et al.,  
136 2003):

$$137 \quad b_{\text{abs}}(\lambda) = b_{\text{ATN}}(\lambda)/C \quad (1)$$

138 As discussed in detail by Corbin et al. (2018), the wavelength-dependence of  $C$  can be expected to be negligible.

139 The loading compensated  $b_{\text{ATN}}$  at  $880$  nm from the AE33 is further used to infer the equivalent-BC mass  
140 concentration,  $M_{\text{eBC}}$ :

$$141 \quad M_{\text{eBC}} = \frac{b_{\text{ATN}}(880 \text{ nm})}{\sigma_{\text{ATN}}(880 \text{ nm})} \quad (2)$$

142 where  $\sigma_{\text{ATN}}$  is the mass attenuation cross-section of BC deposited on the filter of the AE33.  $M_{\text{eBC}}$  inferred from Eq.  
143 (2) only equals the true BC mass concentration,  $M_{\text{BC}}$ , if the applied  $\sigma_{\text{ATN}}$  is identical to the true attenuation cross-  
144 section of BC,  $\sigma_{\text{ATN,BC}}$ , and if light attenuation at  $880$  nm is exclusively due to BC.  $\sigma_{\text{ATN,BC}}(880 \text{ nm})$  can be  
145 inferred from the true MAC of BC,  $\text{MAC}_{\text{BC}}$ , and the true  $C$  value:

$$146 \quad \sigma_{\text{ATN,BC}}(880 \text{ nm}) = \text{MAC}_{\text{BC}}(880 \text{ nm}) * C \quad (3)$$

147 with  $\text{MAC}_{\text{BC}}$  being defined as:

$$148 \quad \text{MAC}_{\text{BC}}(\lambda) = \frac{b_{\text{abs,BC}}(\lambda)}{M_{\text{BC}}} \quad (4)$$

149 where  $b_{\text{abs,BC}}$  is the absorption coefficient due to BC.

150 The manufacturer default values are 1.57 for  $C$  (TFE-coated glass fiber filters) and  $12.2 \text{ m}^2 \text{ g}^{-1}$  for  $\sigma_{\text{ATN}}$  at 880 nm,  
151 which corresponds to a  $\text{MAC}_{\text{BC}}(880 \text{ nm})$  of  $7.77 \text{ m}^2 \text{ g}^{-1}$  at (Gundel et al., 1984, Drinovec et al., 2015). However,  
152 these three parameters depend on aerosol properties. Here, we have determined the  $C$  value by applying Eq. (1) to  
153  $b_{\text{ATN}}$  measured by the aethalometer and the absorption coefficient,  $b_{\text{absMwAA}}$ , measured by a multi-wavelength  
154 absorbance analyser, MWAA (Massabò et al., 2015; Massabò et al., 2013). The  $\text{MAC}_{\text{BC}}(880 \text{ nm})$  was determined  
155 using Eq. (4) to compare  $b_{\text{absMwAA}}$  from the MWAA measurements with EC mass from the Sunset thermo-optical  
156 analyser (see Fig. 1A&B and Section 4.1 for detailed discussion). Following this procedure, the MWAA and Sunset  
157 analyser will be defined as reference methods for absorption coefficient and EC mass concentration, respectively.  
158 Note that data from these reference methods were only available with low time resolution and for a subset of all  
159 samples. Thus, the aethalometer anchored against these reference methods, was used to obtain the wavelength  
160 dependent absorption coefficients and the eBC mass concentrations with high time resolution using Eq. (1) and (2),  
161 respectively. Processing the loading compensated AE33 attenuation coefficients with  $C$  value and  $\text{MAC}_{\text{BC}}$ ,  
162 determined with independent MWAA and Sunset analyser measurements, ensures that the inferred  $b_{\text{abs}}(\lambda)$  (Eq. (1))  
163 and  $M_{\text{eBC}}$  (Eq. (2)) have minimal bias compared to respective true values.

164 **MWAA measurements.** The MWAA (Massabò et al., 2015; Massabò et al., 2013) was used as reference method  
165 for the aerosol absorption coefficient. It measures the absorption coefficient  $b_{\text{absMwAA}}(\lambda)$  of particles deposited on  
166 on standard filter samples. It is composed of five laser diodes, with  $\lambda = 375, 407, 532, 635$  and  $850 \text{ nm}$ , acting as  
167 light sources and placed above the filter, an automated sample-changer, and three low-noise UV-enhanced  
168 photodiodes. ~~The first photodiode is placed behind the filter for the analysis of transmittance measurements, while~~  
169 ~~the other two photodiodes are positioned at specific angles between the sources and the loaded filter to perform~~  
170 ~~reflectance measurements.~~ The first photodiode is placed behind the filter for transmittance measurements ( $0^\circ$   
171 relative to the incident light, 1.5 cm from the sample), while the other two photodiodes are positioned at  $125^\circ$  and  
172  $165^\circ$  (11 cm from the sample) to collect the back scattered light. These transmittance and reflectance measurements  
173 are used together with a radiative transfer model (Hänel et al., 1987), which takes into account multiple scattering  
174 within the particle/filter layer, to retrieve both the total optical thickness and the particle-filter-layer single scattering  
175 albedo, providing the absorption coefficient  $b_{\text{absMwAA}}(\lambda)$  values. These calculations largely follow the approach

176 implemented in the multi-angle absorption photometer (MAAP, Petzold and Schönlinner, 2004).

177 **UV-visible absorbance measurements of extracted aerosols.** Filter samples were extracted for UV-visible  
178 absorbance measurements in 10 mL ultrapure water or methanol in an ultrasonic bath for 20 min at 30 °C. Samples  
179 were subsequently briefly vortexed (1 min) and filtered with 0.45 µm nylon membrane syringe filters following the  
180 procedure described in Daellenbach et al. (2016). Absorption spectra were measured from 280 to 500 nm using a  
181 UV-visible spectrophotometer (Ocean Optics) coupled to a 50-cm long-path detection cell (Krapf et al., 2016). Light  
182 attenuation by the OA in solution,  $ATN_{OA-sol}$ , at a given wavelength was recorded as the logarithm of the ratio of  
183 signal intensities of the reference (solvent) ( $I_0$ ) and the sample ( $I$ ), both corrected for background signals with the  
184 light source off. From  $ATN_{OA-sol}$ , the absorption coefficient of OA in solution,  $b_{abs,OA-sol}(\lambda)$ , can be quantified as:

$$185 \quad b_{abs,OA-sol}(\lambda) = \frac{ATN_{OA-sol}(\lambda)}{l} \quad (5)$$

186 where  $l$  is the optical path length.

187 The absorbance measurements are aimed at inferring the imaginary part of the refractive index. For this,  
188  $b_{abs,OA-sol}(\lambda)$  is transformed to the absorption coefficient of the bulk OA in the pure form,  $b_{abs,OA-bulk}$  (Sun et al.,  
189 2007):

$$190 \quad b_{abs,OA-bulk}(\lambda) = b_{abs,OA-sol}(\lambda) \frac{\rho_{OA}}{\frac{m_{OA}}{V_{solvent}}} \quad (6)$$

191 where  $\rho_{OA}$  is the bulk density of OA (assumed to be 1.5 g cm<sup>-3</sup>, typical of wood-burning OA; (Corbin et al., 2015a;  
192 Moosmüller et al., 2009; Sun et al., 2007)),  $m_{OA}$  is the extracted OA mass, and  $V_{solvent}$  is the solvent volume. The  
193 bulk absorption coefficient directly leads to the imaginary part of the OA refractive index,  $k_{OA}$ , in pure form  
194 (Moosmüller et al., 2009):

$$195 \quad k_{OA}(\lambda) = b_{abs,OA-bulk}(\lambda) \frac{\lambda}{4\pi} \quad (7)$$

196 Inserting Eq. (6) into Eq. (7) eventually provides (Liu et al., 2015a):

$$197 \quad k_{OA}(\lambda) = \frac{\lambda \rho_{OA} V_{solvent}}{4\pi m_{OA}} b_{abs,OA-sol}(\lambda) \quad (8)$$

198 The mass of organics dissolved in the solution could not be quantified. Therefore, we use an upper limit value for  
199  $m_{\text{OA}}$ , approximated as the integral of AMS-measured OA mass concentration times sample flow rate over the filter-  
200 sampling period. Accordingly, the resulting  $k_{\text{OA}}$  values represent lower limits for the true values, as the OA  
201 extraction efficiency was not accounted for. ~~Higher  $k_{\text{OA}}$  values based on online absorption coefficient measurements~~  
202 ~~compared to those calculated based on Equation 8 may be related to low OA extraction efficiency or to non-~~  
203 ~~extractable highly absorbing material and results shall be discussed accordingly.~~ If the OA extraction efficiency was  
204 less than unity, then the absorption (or MAC) predicted from our solvent-extraction measurements would be less  
205 than that measured (or calculated) using our real-time measurements (MWAA-calibrated aethalometer).  
206

### 207 **2.3 Uncertainty analysis**

208 It is important to draw a clear distinction between uncertainties related to measurement precision and accuracy and  
209 those related with experimental variability. In this section we discuss the quantifiable and unquantifiable  
210 uncertainties related with the different measurements. In the result section, we will present our confidence levels on  
211 the average parameters determined based on the experimental variability, which we judge to be the main source of  
212 variance in the data.

213 **Quantifiable uncertainties.** The estimated uncertainty in the AMS-derived OA mass concentrations is ~25%,  
214 which includes both potential biases and precision. This estimate is based on the variation in the AMS calibration  
215 factors and estimated uncertainties in the SMPS used for the AMS calibration (Bruns et al., 2015, 2016).  
216 Uncertainties related to particle transmission efficiency in the AMS are considered negligible for the particles  
217 sampled here (Liu et al., 2007), whose volume size distribution falls within the range transmitted efficiently by the  
218 AMS aerodynamic lens (see Fig. S4). The bounce-related collection efficiency (CE) of the AMS was concluded to  
219 be unity for wood-burning OA in the literature reviewed by Corbin et al. (2015b; in their Section S1.2). For the  
220 present data, the comparison between the SMPS mass (predicted from fitted volume distributions using a density of  
221  $1.5 \text{ g cm}^{-3}$ ) and the total PM predicted as AMS-OA+eBC, suggest a CE value between 0.7 and 1.0 (19% relative  
222 uncertainty), consistent with average literature values and the uncertainties estimates. The uncertainty in EC mass  
223 concentration, estimated from measurement repeats based on the EUSAAR2 protocol only, is within 7% in our case.  
224 The precision uncertainty in the aethalometer attenuation measurements was estimated as  $15 \text{ Mm}^{-1}$  based on the

225 standard deviation of its signals prior to aerosol being injected into the smog chamber. The MWAA data have an  
226 estimated noise level and precision of 12 /Mm and 10% respectively, and these uncertainties have been added in  
227 quadrature to provide the overall uncertainties shown, for example, as error bars in Fig. 1 below. To compare the  
228 MWAA and aethalometer measurements, we determined  $b_{\text{abs,MWAA},880\text{nm}}$  by extrapolating the absorption coefficients  
229 measured at 850 nm to 880 nm using an  $\alpha$ -value determined from the ratio between the absorption coefficients at  
230 850 nm and 635nm. The uncertainty associated with this extrapolation is considered negligible relative to the overall  
231 MWAA uncertainty.

232 **Possible unquantified uncertainties.** There are significant uncertainties in the measurement of aerosol absorption  
233 using filter-based techniques (e.g., Collaud Coen et al., 2010). Here, we have used MWAA measurements as a  
234 reference to scale the aethalometer data, using a single C value. The correction factor C, which accounts for  
235 scattering effects within the filter matrix (Drinovec et al., 2015), may depend on the aerosol sample (Collaud Coen  
236 et al., 2010). In this study, we evaluated the variability in this factor for our primary and aged samples, by directly  
237 comparing the aethalometer to MWAA measurements, as discussed below. The MWAA has been previously  
238 validated against a polar nephelometer and a MAAP (Massabo et al., 2013), which, in turn, has been validated  
239 against numerous in situ methods (e.g., Slowik et al., 2007). The excellent correlation between MWAA and EC in  
240 our study (discussed below) supports the high confidence in the MWAA filter based absorption measurements  
241 conducted here. Another significant source of uncertainty in filter-based absorption measurements is the possible  
242 sorption (or evaporation) of volatile organics on (or from) the filter material. This may lead to an overestimation (or  
243 underestimation) of OA absorption. However, we have minimized sorption artefacts by utilizing a charcoal denuder.  
244 We have obtained an excellent correlation between OA absorption measurements derived from the MWAA-  
245 calibrated aethalometer and from quartz filter samples (see discussion below, Fig. 6 in the main text and S13 in the  
246 supplementary information). Although both of these techniques involved filter sampling, their sampling timescale is  
247 an order of magnitude different, and a difference is therefore expected if sorption (or evaporation) caused a  
248 substantial bias in our results. We therefore conclude that it is unlikely that artifacts associated with filter sampling  
249 have biased the absorption measurements. Finally, uncertainties related to pyrolysis during thermo-optical analysis  
250 may bias EC measurements. Such uncertainties arise from unstable organic compounds, and can be significant for  
251 biomass-burning samples, leading to biases on the order of 20% for EC (e.g. Schauer et al., 2003; Yang and Yu.,  
252 2007). To minimize these biases we applied the EUSAAR2 protocol. The optical properties of such organics are

253 generally different from BC; therefore, the excellent correlation between MWAA and EC data in Fig. 1A suggest  
254 that pyrolysis effects were not a major source of uncertainty in our data set.

255

### 256 3. OPTICAL PROPERTIES ANALYSIS

#### 257 3.1 Determination of absorption Ångström exponents and mass absorption cross-sections

258 In this section we describe the methodology adapted for the determination of the mass absorption cross-sections  
259 (MACs) for the different aerosol material from the Sunset, MWAA and aethalometer measurements. The  
260 assumptions and limitations underlying these calculations are clearly stated. We also explain the relationship  
261 between the MACs and the wavelength dependence of the overall absorption.

262 **Definition of the absorption Ångström exponent .** The wavelength dependence of the overall absorption due to  
263 both BC and BrC has often been described assuming a power law:

$$264 \quad b_{\text{abs}}(\lambda) \propto \lambda^{-\alpha} \quad (9)$$

265 where  $\alpha$  is the Ångström absorption exponent, often determined by fitting the absorption coefficient measurements  
266 across the entire wavelength range. Eq. (9) is an empirical simplification, which breaks down when different  
267 components having different spectral dependence contribute to the absorption, e.g. a mix of BrC and black carbon  
268 (e.g., Moosmüller et al., 2011). In practice, different values of  $\alpha$  would be obtained for different choices of  $\lambda$  ranges,  
269 and therefore we alternatively calculated two-wavelength absorption exponents according to

$$270 \quad \alpha(\lambda, \lambda_{\text{ref}}) = - \frac{\ln\left(\frac{b_{\text{abs}}(\lambda)}{b_{\text{abs}}(\lambda_{\text{ref}})}\right)}{\ln\left(\frac{\lambda}{\lambda_{\text{ref}}}\right)} \quad (10)$$

271 where  $\lambda$  is a wavelength of interest (in nm) and  $\lambda_{\text{ref}}$  is the reference wavelength, here 880 nm. This reference  
272 wavelength was chosen, because BC is expected to fully dominate light absorption in this range (Laskin et al.,  
273 2015).



274 Black carbon is known to have an  $\alpha$  between 0.9 and 1.1 (Bond et al., 2013; Kirchstetter et al., 2004; Liu et al.,  
 275 2015b), whereas BrC, which preferentially absorbs at shorter wavelength, has a higher  $\alpha$  (Laskin et al., 2015; Saleh  
 276 et al., 2013). Thus, we interpret an increase of  $\alpha(\lambda, \lambda_{ref})$  of the total aerosol as due to an increased contribution of  
 277 BrC to the total absorption.  $\alpha(\lambda, \lambda_{ref})$  can potentially change due to other effects such as a wavelength dependent  
 278 lensing effect on absorption by BC (e.g., Lack and Langridge, 2013) or the restructuring of BC aggregates during  
 279 aging. The former effect was negligible under our conditions, as elaborated on below. The latter, if it occurs during  
 280 aging, would be attributed to SOA absorption in our approach. However, this is not an issue if our values are  
 281 accordingly applied in e.g. model simulations, following the same assumption as in our approach. This means that  
 282 the potential restructuring effects must implicitly be considered within the  $MAC(\lambda)$  of SOA, while the  $MAC(\lambda)$  of  
 283 BC must be kept fixed.

### 284 3.2 Determination of $MAC_{BC}$ and $MAC_{POA}$ using the absorption Ångström exponent

285 In a mixture of  $n$  absorbing species, the total absorption at any wavelength may be written as the sum of the  
 286 absorbance of each of the species. Accordingly, Eq. (10) can be expressed for a multi-component system

$$287 \alpha(\lambda, \lambda_{ref}) = \frac{1}{\ln(\lambda_{ref}/\lambda)} \ln \left( \frac{\sum_{i=1}^n b_{abs,i}(\lambda)}{\sum_{i=1}^n b_{abs,i}(\lambda_{ref})} \right) = \frac{1}{\ln(\lambda_{ref}/\lambda)} \ln \left( \frac{\sum_{i=1}^n M_i MAC_i(\lambda)}{\sum_{i=1}^n M_i MAC_i(\lambda_{ref})} \right) \quad (11)$$

288 where the right hand side follows the general definition of MAC along the lines of Eq. (4).  $M_i$  and  $MAC_i$  are the  
 289 mass concentration and MAC, respectively, of the  $i^{th}$  species, with  $n$  absorbing species in total. By considering that  
 290 the light absorption at  $\lambda_{ref} = 880$  nm is exclusively due to BC, and by defining BC to be the  $n^{th}$  species, Eq. (11)  
 291 can be written as

$$292 \alpha(\lambda, 880nm) = \frac{1}{\ln(880nm/\lambda)} \ln \left( \frac{MAC_{BC}(\lambda)}{MAC_{BC}(880nm)} + \sum_{i=1}^{n-1} \frac{M_i MAC_i(\lambda)}{b_{abs}(880nm)} \right) \quad (12)$$

293 In Eq. (12), the summation now only goes over the  $n-1$  organic species, which contribute to light absorption.  
 294 The fresh combustion aerosol exclusively contains BC and POA as absorbing species. For the data at time  $t_0$  before  
 295 the start of photo-oxidative aging, Eq. (12) simplifies to:

$$296 \alpha(t_0, \lambda, 880nm) = \alpha_{BC+POA}(t_0, \lambda, 880nm)$$

$$297 \quad = \frac{1}{\ln(880\text{nm}/\lambda)} \ln \left( \frac{\text{MAC}_{\text{BC}}(t_0, \lambda)}{\text{MAC}_{\text{BC}}(t_0, 880\text{nm})} + \frac{M_{\text{OA}}(t_0) \text{MAC}_{\text{POA}}(t_0, \lambda)}{b_{\text{abs}}(t_0, 880\text{nm})} \right) \quad (13)$$

298 In Eq. (13),  $M_{\text{OA}}(t_0)$  is the mass concentration of primary organic aerosol measured by the AMS at  $t_0$ .  
 299  $\text{MAC}_{\text{BC}}(t_0, 880\text{nm})$  was inferred from the MWAA and Sunset thermo-optical analysis and shown to be independent  
 300 of the experimental conditions (Section 4.1; Fig. 1A). Absorption coefficients  $b_{\text{abs}}(t_0, \lambda)$  are obtained from the high  
 301 time resolution attenuation measurements by the aethalometer referenced to the MWAA absorption measurements  
 302 as described above.  $\alpha(t_0, \lambda, 880 \text{ nm})$  is derived from  $b_{\text{abs}}(t_0, \lambda)$  and  $b_{\text{abs}}(t_0, 880 \text{ nm})$  using Eq. (10). We have  
 303 intentionally formulated of Eq. (13) as such to highlight that the retrieved  $\text{MAC}_{\text{OA}}(t, \lambda)$  depends mainly on the input  
 304  $M_{\text{OA}}$ . Correspondingly, the retrieved  $\text{MAC}_{\text{OA}}(t, \lambda)$  is mainly sensitive to potential AMS calibration biases. This leaves  
 305 only 2 free parameters in Eq. (13),  $\text{MAC}_{\text{BC}}(t_0, \lambda)$  and  $\text{MAC}_{\text{POA}}(t_0, \lambda)$ . These were determined by fitting Eq. (13) to  
 306  $\alpha(t_0, \lambda, 880 \text{ nm})$ ,  $M_{\text{OA}}(t_0)$ ,  $\text{MAC}_{\text{BC}}(t_0, 880\text{nm})$  and  $b_{\text{abs}}(t_0, 880\text{nm})$  data measured in all experiments for fresh  
 307 emissions at  $t_0$ . This approach contains the implicit assumption that the two MAC values are also independent of  
 308 experimental conditions, and therefore these MACs should be considered as average values. The accuracy of these  
 309 MAC values obviously depends on the accuracy of the absorption and mass measurements. First, a systematic bias  
 310 in the  $C$  value potentially caused by a systematic bias in the MWAA measurements propagates to an identical bias in  
 311 both  $\text{MAC}_{\text{BC}}(t_0, \lambda)$  and  $\text{MAC}_{\text{POA}}(t_0, \lambda)$ . Second, a systematic bias in the Sunset EC mass measurements yields a  
 312 corresponding inverse bias in  $\text{MAC}_{\text{BC}}(t_0, \lambda)$ , while  $\text{MAC}_{\text{POA}}(t_0, \lambda)$  remains unaffected. Third, a systematic bias in the  
 313 AMS POA mass yields a corresponding inverse bias in  $\text{MAC}_{\text{POA}}(t_0, \lambda)$ , while  $\text{MAC}_{\text{BC}}(t_0, \lambda)$  remains unaffected. Eq.  
 314 (13) shows that  $\alpha$  of the primary aerosol at a certain wavelength is largely driven by  $\text{MAC}_{\text{POA}}(t_0, \lambda)$ , i.e. the optical  
 315 properties of POA, and by the ratio  $\frac{M_{\text{OA}}(t_0)}{b_{\text{abs}}(t_0, 880\text{nm})}$ , which reflects the relative contributions of POA and BC to total  
 316 primary aerosol mass.

### 317 3.3 Determination of $\text{MAC}_{\text{SOA}}$

318 The MAC of SOA,  $\text{MAC}_{\text{SOA}}$ , can be generally defined as:

$$319 \quad \text{MAC}_{\text{SOA}} = \frac{b_{\text{abs,SOA}}}{M_{\text{SOA}}} \quad (14)$$

320 where  $b_{\text{abs,SOA}}$  and  $M_{\text{SOA}}$  are the absorption coefficient and mass concentration of SOA, respectively. In the aged  
 321 aerosol, which contains the absorbing species BC, POA and SOA,  $b_{\text{abs,SOA}}$  is the difference of the total absorption  
 322 minus the absorption by POA and BC:

323  $b_{\text{abs,SOA}}(t, \lambda) = b_{\text{abs}}(t, \lambda) - b_{\text{abs,POA+BC}}(t, \lambda)$  (15)

324 The absorption by POA and BC in the aged aerosol is a priori unknown, but can be calculated under certain  
 325 assumptions. The first assumption is that SOA does not contribute to absorption at 880 nm:  
 326  $b_{\text{abs,POA+BC}}(t, 880 \text{ nm}) \equiv b_{\text{abs}}(t, 880 \text{ nm})$ . The second assumption is that the two- $\lambda$   $\alpha$  values of primary emissions  
 327 do not change during aging  $\alpha_{\text{POA+BC}}(t, \lambda, 880 \text{ nm}) \equiv \alpha_{\text{POA+BC}}(t_0, \lambda, 880 \text{ nm})$ . The latter approximation is based on  
 328 the underlying assumptions that the MAC of POA is not altered by aging and that the proportions of POA and BC  
 329 mass lost to the wall are identical. Under these assumptions  $b_{\text{abs,POA+BC}}$  becomes:

330  $b_{\text{abs,POA+BC}}(t, \lambda) = b_{\text{abs}}(t, 880 \text{ nm}) \left( \frac{880 \text{ nm}}{\lambda} \right)^{\alpha_{\text{POA+BC}}(t_0, \lambda, 880 \text{ nm})}$  (16)

331 Note that inferring  $b_{\text{abs,POA+BC}}(t, \lambda)$  from  $b_{\text{abs}}(t, 880 \text{ nm})$  implicitly accounts for the decrease in the BC and POA  
 332 absorption due to wall losses.

333  $M_{\text{SOA}}$  was obtained as total organic minus POA mass concentration:

334  $M_{\text{SOA}}(t) = M_{\text{OA}}(t) - M_{\text{POA}}(t)$  (17)

335 The POA mass concentration in the aged aerosol can be inferred from the initial OA mass concentration in the fresh  
 336 emissions by accounting for the wall losses using Eq. (S1) and the wall loss time constant  $\tau$  (see Section Wall loss  
 337 corrections in the SI):

338  $M_{\text{POA}}(t) = M_{\text{OA}}(t_0) \exp(\tau^{-1}t)$  (18)

339 Inserting Eq. (15) - (18) into Eq. (14) provides the final equation for inferring  $\text{MAC}_{\text{SOA}}$ .

340  $\text{MAC}_{\text{SOA}}(t, \lambda) = \frac{b_{\text{abs}}(t, \lambda) - b_{\text{abs}}(t, 880 \text{ nm}) \left( \frac{880 \text{ nm}}{\lambda} \right)^{\alpha_{\text{POA+BC}}(t_0, \lambda, 880 \text{ nm})}}{M_{\text{OA}}(t) - M_{\text{OA}}(t_0) \exp(\tau^{-1}t)}$  (19)

341  
 342  $\text{MAC}_{\text{SOA}}$  can be calculated for every data point in time and for all aethalometer wavelengths from 370 to 660 nm  
 343 ( $\text{MAC}_{\text{SOA}}$  defined to be zero at  $\lambda \geq 880 \text{ nm}$ ), as all quantities on the right hand side of Eq. (19) are available from  
 344 either the aethalometer or AMS measurements or are otherwise known. It can be seen from Eq. (19) that the mass  
 345 concentrations used to calculate  $\text{MAC}_{\text{SOA}}$  solely originate from AMS data, thus being consistent with the calculation  
 346 of  $\text{MAC}_{\text{POA}}$  (see above). Eq. (19) is based on the assumption that POA is “chemically inert”, i.e. no chemically

347 induced changes of  $M_{\text{POA}}$  and  $\text{MAC}_{\text{POA}}$  occur. ~~Such chemically induced changes of absorption by POA, if they~~  
 348 ~~occur, are assigned to the absorption by SOA, thus resulting in a corresponding adjustment of the inferred  $\text{MAC}_{\text{SOA}}$ .~~  
 349 Such chemically induced changes of absorption coefficient by POA, through a change of  $M_{\text{POA}}$  or  $\text{MAC}_{\text{POA}}$ , if they  
 350 occur, are assigned to the absorption by SOA, thus resulting in a corresponding adjustment of the inferred  $\text{MAC}_{\text{SOA}}$ .

### 351 3.4 Mie calculation to relate $k_{\text{OA}}$ with $\text{MAC}_{\text{OA}}$

352 The imaginary part of the refractive index of an aerosol component is an intensive material property. However, the  
 353 MAC of such an aerosol component additionally depends on the size and morphology of the aerosol (except for the  
 354 Rayleigh regime). The online aerosol absorption measurements provide estimates for MAC values, while the UV-  
 355 visible absorbance measurements of filter extracts provide the imaginary part of the refractive index. We used Mie  
 356 calculations in order to compare the two quantities. The  $k_{\text{OA}}(\lambda)$  obtained from the filter extracts is converted to a  
 357  $\text{MAC}_{\text{OA,bulk}}$  by assuming that all OA is present in homogeneous spherical particles with a diameter distribution  
 358 identical to the mobility diameter distribution measured by the SMPS. In this manner,  $\text{MAC}_{\text{OA,bulk}}$  becomes equal to  
 359 the mass-weighted average (=volume-weighted average) of the diameter dependent MAC:

$$360 \text{MAC}_{\text{OA,bulk}}(\lambda, n_{\text{OA}}, k_{\text{OA}}, \rho_{\text{OA}}) = \frac{\sum_i N_i d_i^3 \text{MAC}_i^{\text{Mie}}(\lambda, n_{\text{OA}}, k_{\text{OA}}, \rho_{\text{OA}})}{\sum_i N_i d_i^3} \quad (20)$$

361 Here,  $N_i$  and  $d_i$  are the number of particles and particle diameter, respectively, in the  $i^{\text{th}}$  size bin, and  $n_{\text{OA}}$  is the real  
 362 part of the refractive index of the OA (which is assumed to be  $n_{\text{OA}} = 1.5$  typical for organic material; Lu et al.,  
 363 2015). The MAC of particles with diameter  $d_i$ ,  $\text{MAC}_i^{\text{Mie}}$ , was calculated using the Mie Code by Peña and Pal (2009)  
 364 (incorporated into Igor Pro 6.3, WaveMetrics, OR, USA by Taylor et al., 2015).  $\text{MAC}_i^{\text{Mie}}$  also depends on the density  
 365 of OA, for which we assume a value of  $\rho_{\text{OA}} = 1.5 \text{ g cm}^{-3}$  (see Section 2.2), as the volume specific absorption cross-  
 366 section obtained from Mie theory needs to be converted to a mass specific absorption cross-section. We note that as  
 367 we have used the same value of  $\rho_{\text{OA}}$  in the calculation of both  $\text{MAC}_i^{\text{Mie}}$  and  $k_{\text{OA}}(\lambda)$ ,  $\text{MAC}_{\text{OA,bulk}}$  becomes  
 368 independent of the assumed  $\rho_{\text{OA}}$  value.

369 Assuming spherical particles and neglecting the presence of BC in these particles may seem inappropriate. However,  
 370 calculations considering BC and assuming core-shell morphology revealed (1) limited sensitivity of the resulting  
 371  $\text{MAC}_{\text{OA}}$  to this assumption and (2) a higher than measured lensing effect. Therefore, a substantial fraction of the OA  
 372 seems to be externally mixed and to dominate the measured size distribution (see also Section 4.1). ~~The uncertainty~~

373 in the  $MAC_{OA}$  inferred from  $k_{OA}$  of the UV-visible absorbance measurements was estimated by combining an  
374 estimated 20 % precision with a detection limit of  $0.3 \text{ m}^2 \text{ g}^{-1}$  in quadrature.

375

## 376 4. RESULTS AND DISCUSSION

### 377 4.1 Verification of $MAC_{BC}$ and $C$ value

378 As mentioned above, the determination of  $MAC_{BC}(880\text{nm})$  requires the determination of the absorption coefficients  
379 at  $\lambda$  and the BC mass. We used the aethalometer to obtain the absorption coefficients with high time resolution,  
380 while absolute values were scaled to match MWA data, which we defined as our reference method. The  
381 aethalometer was also used to obtain eBC mass concentrations with high time resolution, while absolute values were  
382 scaled to match EC mass measured by the Sunset thermo-optical measurement using the EUSAAR 2 protocol,  
383 which we defined as our reference method. Here we start by proving the concept of our scaling approaches and  
384 provide average values for  $MAC_{BC}(880\text{nm})$  and aethalometer  $C$  which are required subsequently.

385 Figure 1A shows the correlation between the MWA measured absorption coefficient at 880 nm and the Sunset  
386 thermo-optical EC mass measurements. MWA absorption measurements at 880 nm is determined by  
387 extrapolating the absorption coefficients at 850 nm using an  $\alpha$  determined from the ratio between the  
388 absorption coefficients at 850 nm and 635nm. The corresponding  $MAC_{BC}(880\text{nm})$ , determined as the slope of  
389 the linear fit through all data, is  $4.6 \pm 0.7 \text{ m}^2 \text{ g}^{-1}$ . This value matches the data at all three levels of aging, i.e. for the  
390 primary, Aged1 and Aged2 filter samples, within experimental uncertainty (see Figure S2 in the Supplement for  
391 more information). This average  $MAC_{BC}(880\text{nm})$  is also very similar to values reported for “pure” BC ( $4.7 \pm 0.7$   
392  $\text{m}^2 \text{ g}^{-1}$  at 880 nm) (Bond et al., 2006), indicating no significant lensing effect on absorption by BC from primary or  
393 secondary OA. This can also be observed from the time-resolved attenuation measurements by the aethalometer at  
394 880 nm (Figure S3), suggesting that little (<10%) to no increase in the attenuation coefficients upon SOA formation.  
395 If the OA and the BC were internally mixed, the observed variability in the mass fraction of OA ( $f_{OA}$ ) from 0.1 to  
396 0.9 for the fresh and aged samples would result in a high variability in the  $MAC_{BC}(880\text{nm})$ , with values higher than  
397 those reported in the literature, according to Mie calculations assuming core-shell internal mixtures. However, this is  
398 not the case. Based on this observation, we conclude that the particles studied are likely not core-shell internal

399 mixtures, although we have measured a mono-modal aerosol population growing during SOA production (Figure  
400 S4). An explanation for the occurrence of an external mixture could be that the primary OA and BC particles may  
401 have been externally mixed after these species were emitted separately during combustion, preferentially during the  
402 pre-ignition and flaming phases, respectively (Corbin et al., 2015a, 2015b; Heringa et al., 2011). These phases may  
403 occur consecutively during a burn or simultaneously in different parts of the stove.  $MAC_{BC}(880nm)$  found to be  
404 constant supports our approach described in Section 2.2 using scaled aethalometer data for BC mass and treating  
405  $MAC_{BC}$  at all other wavelengths as a constant across all experiments during the data retrieval process described in  
406 Section 3.1.

407 Figure 1B shows the correlation between  $b_{ATN,AE33}$  and  $b_{abs,MWAA}$  measured by the aethalometer and MWAA,  
408 respectively. The two variables correlated very well, indicating a constant aethalometer  $C$  value, which is the ratio  
409 between  $b_{ATN,AE33}$  and  $b_{abs,MWAA}$  (Equation 1), of  $3.0 \pm 0.3$ , independent of the type of the aerosol sampled. This is  
410 also reflected in the probability density function of individual  $C$  values shown in Figure S2 where the standard  
411 deviation is found to be as small as  $\sigma_C \sim \pm 10\%$ . Such constant ratio justifies our approach of applying this single  
412  $C$  value for all conditions in order to scale the time-resolved attenuation measurements by the aethalometer to the  
413 MWAA reference method.

414 Note, the manufacturer's default values, which were not applied in our case, are 1.57 for  $C$  and  $12.2 \text{ m}^2 \text{ g}^{-1}$  for  
415  $\sigma_{ATN}$  at 880 nm, which implies an underlying  $MAC_{BC}(880 \text{ nm})$  of  $7.77 \text{ m}^2 \text{ g}^{-1}$  (Gundel et al., 1984, Drinovec et al.,  
416 2015). Therefore, factory default  $b_{abs}(\lambda)$  would have a substantial systematic high bias for the wood combustion  
417 aerosols of this study. Meanwhile, the  $\sigma_{ATN}$  calculated at 880 nm, which is the product of the  $C$  value and  $MAC_{BC}$   
418 (Equation 3), is consistent with the manufacturer value of  $\sigma_{ATN}$  ( $\sigma_{ATN}$  values determined here are 15% higher,  
419  $13.8 \text{ m}^2 \text{ g}^{-1}$  in this study compared to the value of  $12.2 \text{ m}^2 \text{ g}^{-1}$  provided by the manufacturer), and the factory default  
420  $M_{eBC}$  would agree well with the true  $M_{BC}$ , determined here.

421 We have independently determined the  $MAC_{BC}(880nm)$  and the aethalometer  $C$  values under our conditions, as  
422 follows. We determined  $MAC_{BC}(880nm)$  from the regression between the absorption coefficients at 880 nm  
423 obtained from the MWAA and the EC mass measured by the Sunset analyzer (Fig. 1A). The slope of this regression  
424 may be used to estimate the  $MAC_{BC}(880nm)$ , which we retrieved as  $4.7 \pm 0.3 \text{ m}^2 \text{ g}^{-1}$  by an uncertainty-weighted  
425 linear least-squares fit. The corresponding intercept was not significantly different from zero ( $-3 \pm 3 / \text{Mm}$ ). Our  
426  $MAC_{BC}(880nm)$  is not statistically significantly different from the value recommended by Bond et al., (2006) for

427 externally-mixed BC (extrapolating their  $MAC_{BC}(550nm)$  to 880 nm by assuming  $\alpha=1$  provides  $MAC_{BC}(880nm)=$   
428  $4.7 \pm 0.7 m^2 g^{-1}$ ). The strong correlation between  $b_{abs,MWAA,880nm}$  and EC in Fig. 1A shows that  $MAC_{BC}(880nm)$  did  
429 not vary with aging during our study (see also Fig. S2-a). It also indicates that measurement artefacts for both  
430 instruments were negligible, as the fundamental differences between the two techniques mean that any artefacts are  
431 unlikely to be similar between them (charring for EC vs. adsorption artefacts for MWAA). Our absorption  
432 coefficient measurements also provide insights into particle mixing state in this study. Since a single MAC  
433 adequately described our samples at all levels of aging (Fig. 1A and Fig. S2-a), in spite of a factor of 3.3 average  
434 increase in the aerosol mass, our samples cannot be adequately described by a core-shell Mie model. Such a core-  
435 shell model would predict an absorption enhancement by a factor of  $\sim 1.8$  (Bond et al., 2006) for the observed OA  
436 mass increase with aging, which was not observed in our case. This observation is also supported by the time  
437 resolved attenuation measurements at 880 nm using the aethalometer (Fig. S3), suggesting that little ( $<10\%$ ) to no  
438 increase in the attenuation coefficients upon SOA formation. We emphasize that this conclusion does not indicate  
439 that no internal mixing occurred, but rather that the simplified concept of negligible mixing better describes our data  
440 than the equally simplified concept of a core-shell description of coatings that completely envelop the central BC  
441 core. This may be due to the complex morphology of internally-mixed BC, which has been previously observed for  
442 wood burning particles (e.g., China et al., 2013; Liu et al., 2015; Liu et al., 2017). It may also be related to the fact  
443 that OA and BC are emitted during separate phases of combustion. OA rich particles are emitted during the pre-  
444 flaming pyrolysis stage of combustion, whereas most BC is emitted during flaming combustion (Corbin et al.,  
445 2015a, 2015b; Haslett et al., 2018; Heringa et al., 2011). These two stages of combustion may coexist in different  
446 regions of the stove, particularly during simulated real-world usage. As lensing effect was negligible in our case, we  
447 have assumed that the aerosol optically behaves as an external mix between BC and BrC during Mie calculation (see  
448 section 3.4). We note that while this assumption is important for estimating the BC absorption, the conclusion drawn  
449 about the BrC absorption are not very sensitive to the assumed morphology.

450 We determined time-resolved wavelength-dependent absorption coefficients as follows. We used the aethalometer to  
451 obtain filter attenuation coefficients with high time resolution, which were then calibrated to obtain absorption  
452 coefficients by deriving the factor C (Eq. (1)) using the MWAA measurements of filter samples. C was obtained  
453 from an uncertainty-weighted linear least-squares fit as  $3.0 \pm 0.2$  (Fig. 1B); the intercept of the fit was not  
454 significantly different from zero, within two standard deviations ( $-17 \pm 14$ ). A very strong correlation could be  
455 observed between MWAA and aethalometer (Fig. 1B), implying that C is independent of the type of the aerosol

456 sampled (see also Fig. S2-B). Therefore, we used a single C value to obtain time-resolved wavelength-dependent  
457 absorption coefficients from the aethalometer attenuation measurements at the different wavelengths for primary and  
458 aged aerosols.

459 Note that the manufacturer's default values, which were not applied in our case, are 1.57 for C (using TFE-coated  
460 glass fiber filters) and  $12.2 \text{ m}^2 \text{ g}^{-1}$  for  $\sigma_{\text{ATN}}$  at 880 nm (Gundel et al., 1984, Drinovec et al., 2015). The C value  
461 determined here is larger than the manufacturer-default value for the AE33, resulting in smaller absorption  
462 coefficients. However, the calculated  $\sigma_{\text{ATN}}$  at 880 nm ( $13.8 \text{ m}^2 \text{ g}^{-1}$ ), which can be retrieved as the product of the C  
463 value and  $\text{MAC}_{\text{BC}}(880\text{nm})$  (Eq. (3)), is similar to the factory-default  $\sigma_{\text{ATN}}$ . Therefore, our calibrated  $M_{\text{eBC}}$ , calculated  
464 from the attenuation coefficients using  $\sigma_{\text{ATN}}$  (Eq. (2)), are similar to the factory-default  $M_{\text{eBC}}$ . We note that  $M_{\text{eBC}}$  has  
465 not been used for  $\text{MAC}_{\text{OA}}$  calculations, and is only used for the calculation of the mass fractions of BC and OA for  
466 display purposes (Fig. 2, 3, 7 and 8).

#### 467 4.2 Optical properties of BC, POA, and SOA

468 In this section we derive the wavelength dependent mass absorption cross-sections for BC, POA and SOA. In Fig. 2,  
469 we display the evolution of  $\alpha(370\text{nm}, 880\text{nm})$  as a function of OH exposure. Fig. 3 shows the relationship between  
470  $\alpha(\lambda, 880\text{nm})$  and  $f_{\text{OA}}$  for primary and aged aerosols.

471  **$\alpha$  of primary emissions.** The  $\alpha(370\text{nm}, 880\text{nm})$  values computed for the primary aerosol (OH exposure = 0  
472 molecules  $\text{cm}^{-3} \text{ h}$ ) ranged between 1.3 and 1.7 (Fig. S5), which is within the range reported previously for biomass-  
473 burning emissions (Kirchstetter et al., 2004; Lewis et al., 2008; Zotter et al., 2016). The  $\alpha(\lambda, 880\text{nm})$  is slightly  
474 higher than that of pure BC (~1.2; Bond et al., 2013; Zotter et al., 2017) for small  $f_{\text{POA}}$ , while increasing  $f_{\text{POA}}$   
475 corresponded to a distinct increase of  $\alpha(\lambda, 880\text{nm})$ . The  $\alpha(\lambda, 880\text{nm})$  is close to that of pure BC (~0.9-1.1; Bond et  
476 al., 2013; Zotter et al., 2017) for small  $f_{\text{POA}}$ , while increasing  $f_{\text{POA}}$  corresponded to a distinct increase of  
477  $\alpha(\lambda, 880\text{nm})$ . This increase provides clear evidence for the contribution of primary BrC to the absorption at lower  
478 wavelengths (shown explicitly in Eq. (13)). The  $f_{\text{POA}}$  ranges from 0.23 to 0.59, which is lower than  $f_{\text{POA}}$  reported for  
479 open burning emissions (e.g.,  $f_{\text{POA}} \sim 0.75$ , Ulevicius et al (2016)), because our wood stove emissions feature a more  
480 efficient combustion. The  $f_{\text{POA}}$  ranges from 0.12 to 0.63, which is lower than  $f_{\text{POA}}$  reported for open burning emissions  
481 (e.g.,  $f_{\text{POA}} \sim 0.75$ , Ulevicius et al (2016)), because our wood-stove emissions feature a more efficient combustion. The  
482 systematic decrease in  $\alpha(\lambda, 880\text{nm})$  with increasing  $\lambda$  reflects the more efficient light absorption by BrC at shorter  
483 wavelengths (Moosmüller et al., 2011), and shows that the power law wavelength dependence is an inaccurate



484 oversimplification for this mixed aerosol. As illustrated in Fig. S5, the observed absorption spectra have steeper  
485 gradients with decreasing wavelength compared to the lines of constant  $\alpha$ . Such systematic increase in  $\alpha(\lambda, 880\text{nm})$   
486 with decreasing  $\lambda$  reflects the more-efficient light absorption by BrC at shorter wavelengths (Moosmüller et al.,  
487 2011), and shows that the power law wavelength dependence is an inaccurate oversimplification for this mixed  
488 aerosol.

489 **Evolution of  $\alpha$  with aging.** Fig. 3B shows that upon aging, the OA fraction rapidly increased (a typical time series  
490 of raw data is shown in Fig. S1), reaching an average value of 0.81 (full range for aged OA:  $0.74 < f_{OA} < 0.89$ ) at  
491 high OH exposures ( $> 2 \times 10^7$  molecules  $\text{cm}^{-3}$  h), and resulting in a corresponding increase of  
492  $\alpha_{\text{BC+POA+SOA}}(370\text{nm}, 880\text{nm})$ . The increase of  $\alpha_{\text{BC+POA+SOA}}(370\text{nm}, 880\text{nm})$  and  $f_{OA}$  were always correlated  
493 and plateaued at OH exposures beyond  $\sim 2 \times 10^7$  molecules  $\text{cm}^{-3}$  h, as seen in Fig. 2. Also, note in Fig. 2 that the  
494 highest  $\alpha_{\text{BC+POA+SOA}}(370\text{nm}, 880\text{nm})$  were reached, on average 1.8, during experiments where the  $f_{OA}$  was highest  
495 Also, note in Fig. 2 that at highest OH exposures, the highest  $\alpha_{\text{BC+POA+SOA}}(370\text{nm}, 880\text{nm})$  were reached, on average  
496 1.8, during experiments where the  $f_{OA}$  was highest. Such strong correlation between SOA formation and  
497  $\alpha_{\text{BC+POA+SOA}}(370\text{nm}, 880\text{nm})$  suggests the production of substantial amounts of brown SOA. A similar  
498 relationship is observed between  $\alpha_{\text{BC+POA+SOA}}(\lambda, 880\text{nm})$  and  $f_{OA}$  for higher wavelengths as shown in Fig. S6.  
499 Similar to the case of POA, a systematic decrease in  $\alpha(\lambda, 880\text{nm})$  with increasing  $\lambda$  is observed, reflecting the  
500 preferential absorption of BrC SOA at shorter wavelengths. We note that  $\alpha_{\text{BC+POA+SOA}}(370\text{nm}, 880\text{nm})$  as a  
501 function of  $f_{OA}$  for all experiments lies below the overall trend for the primary aerosol (dashed line in Fig. 3B),  
502 implying that  $\text{MAC}_{\text{SOA}}(370\text{nm})$  was smaller than  $\text{MAC}_{\text{POA}}(370\text{nm})$ .

503 **Determination of  $\text{MAC}_{\text{BC}}$  and  $\text{MAC}_{\text{POA}}$ .** We determined best-fit values for  $\text{MAC}_{\text{BC}}(\lambda)$  and  $\text{MAC}_{\text{POA}}(\lambda)$  from the  
504 data shown in Fig. 3A. Fig. 3A includes least-squares fits of Eq. (13) to the data, with  $\text{MAC}_{\text{BC}}(\lambda)$  and  $\text{MAC}_{\text{POA}}(\lambda)$  as  
505 fit parameters. The fit results are shown in Table 1. The obtained fit value of  $\text{MAC}_{\text{BC}}(370\text{nm})$  was  $13.7 \text{ m}^2 \text{ g}^{-1}$  (GSD  
506 1.1), higher but not statistically significantly different from the value suggested by Bond et al. (2013) of  $11.1 \text{ m}^2 \text{ g}^{-1}$   
507 with a 95% confidence interval of  $3.5 \text{ m}^2 \text{ g}^{-1}$ , considering  $\alpha_{\text{BC}}=1$ . The obtained fit value of  $\text{MAC}_{\text{BC}}(370\text{nm})$  was  $13.7$   
508  $\text{m}^2 \text{ g}^{-1}$  (GSD 1.1, one-sigma uncertainty 12.4—15.1  $\text{m}^2/\text{g}$ ), higher but not statistically significantly different from the  
509 range estimated based on Bond et al. (2013), considering the uncertainties on both the  $\alpha_{\text{BC}}$  values and the  
510  $\text{MAC}_{\text{BC}}(520\text{nm})$ . Meanwhile, the mean  $\text{MAC}_{\text{POA}}(370\text{nm})$  value, equal to  $5.5 \text{ m}^2 \text{ g}^{-1}$ , obtained under our conditions  
511 for domestic wood burning is  $\sim 2.4$  times higher than that obtained by Saleh et al. (2014) for open biomass burning

512 primary emissions, suggesting the presence of more-strongly absorbing organic material under our conditions (this  
513 comparison is continued in Section 4.3).

514 **Determination of  $MAC_{SOA}$ .** The  $MAC_{SOA}(\lambda)$  values, determined using Eq. (19), are shown in Fig. 4 and Table 1.  
515  $MAC_{SOA}(370\text{nm})$  was  $2.2 \text{ m}^2 \text{ g}^{-1}$  (GSD 1.39), a factor of 2.5 smaller than  $MAC_{POA}(370\text{nm})$ , but approximately an  
516 order of magnitude higher than values reported for ambient oxygenated aerosols or laboratory SOA from biogenic  
517 and traditional anthropogenic precursors such as terpenes and methyl-benzenes (Clarke et al., 2007; Lambe et al.,  
518 2013; Liu et al., 2016; Romonosky et al., 2015). The predominant SOA precursors identified in wood smoke  
519 comprise (methyl)naphthalene(s) and phenol derivatives from lignin pyrolysis (Bruns et al., 2016; Ciarelli et al.,  
520 2016), the oxidation products of which are expected to be highly light absorbing due to the presence of aromatic  
521 moieties in the SOA (Bruns et al., 2016; Laskin et al., 2015). In this regard, it is not surprising that the  
522  $MAC_{SOA}(370\text{nm})$  values obtained here are similarly high as those obtained from methanol-extracted SOA from  
523 guaiacol and naphthalene oxidation ( $0.5\text{--}3.0 \text{ m}^2 \text{ g}^{-1}$ , Romonosky et al., 2015).

524 **Uncertainties and variability in  $MAC_{BC}$ ,  $MAC_{POA}$  and  $MAC_{SOA}$ .** Table 1 shows that the uncertainties in the fitted  
525  $MAC_{BC}(\lambda)$  are relatively low ( $< 10\%$ ), increasing with decreasing  $\lambda$ . By contrast, the uncertainties in the fitted  
526  $MAC_{POA}$  are much higher (GSD = 1.2–1.5) and increase with increasing  $\lambda$ . The relative residuals between the  
527 measured and fitted  $\alpha(\lambda, 880\text{nm})$  for primary emissions showed small biases of only 0.07 (Fig. S7). The  
528 corresponding RMSE (root mean square error) was 0.13, showing that the obtained average values may represent  
529 the data well.  $MAC_{SOA}$  values determined were highly variable between experiments with a GSD = 1.39 and 2.42  
530 for  $\lambda=370 \text{ nm}$  and  $660 \text{ nm}$ , respectively. We expect the variabilities in  $MAC_{SOA}$  and of  $MAC_{POA}$  to be related to  
531 changes in the organic aerosol chemical composition between different burns, since the variability of  $MAC_{BC}(\lambda)$  was  
532 relatively small. In Section 4.3, we discuss this variability further using the results of an additional and independent  
533 analysis.

534 Table 1 shows the fitting errors related with  $MAC_{BC}(\lambda)$ ,  $MAC_{POA}(\lambda)$  and  $MAC_{SOA}(\lambda)$ , arising from our measurement  
535 precision and experimental variability. These fitting errors are greater than our estimated uncertainties in the  
536 absorption coefficients measured by MWAA (10%), and comparable to our estimated uncertainty in OA mass  
537 measured by AMS (30%). The residuals in the fitted  $MAC_{BC}(\lambda)$  are relatively low ( $< 10\%$ ), increasing with  
538 decreasing  $\lambda$ . By contrast, the uncertainties in the fitted  $MAC_{POA}(\lambda)$  are much higher (GSD = 1.2–1.5) and increase  
539 with increasing  $\lambda$ . The relative residuals between the measured and fitted  $\alpha(\lambda, 880\text{nm})$  for primary emissions showed

540 a mean bias and RMSE of 0.07 and 0.13, respectively (Fig. S7), indicating that our fitted MAC results provide a  
541 good description of the data set.  $MAC_{SOA}(\lambda)$  values determined were highly variable between experiments with a  
542 GSD = 1.39 and 2.42 for  $\lambda=370$  nm and 660 nm, respectively. In Fig. S10, we show the distribution of  $MAC_{SOA}(\lambda)$   
543 values as box and whiskers against OH exposure, showing no particular dependence of these values with aging as it  
544 will be discussed below. Therefore, we expect the fitting errors in  $MAC_{SOA}$  and of  $MAC_{POA}$  to be mainly related to  
545 true changes in the organic aerosol chemical composition between different burns, since the variability of  $MAC_{BC}(\lambda)$   
546 was relatively small. In Section 4.3, we discuss this variability further using the results of an additional and  
547 independent analysis.

548  **$MAC_{BC}$ ,  $MAC_{POA}$  and  $MAC_{SOA}$  wavelength dependence.** The relationships between the  $MAC_{SOA}(\lambda)$ ,  $MAC_{POA}(\lambda)$   
549 and  $MAC_{BC}(\lambda)$  and wavelength appear to fall on three unique lines in the range 660 nm to 370 nm when plotted in  
550 log-log space, as shown in Fig. 4 (Fig. S8 shows the same data plotted on a linear scale). This indicates that a power-  
551 law approximation provides a good description of the behavior of individual components within this wavelength  
552 range from 370 nm to 660 nm. Accordingly we fitted the power law coefficients to the data shown in Fig. 4  
553 ( $\ln(MAC_i) = \ln(A_i) + \alpha_i \ln(\lambda)$ , with  $i = BC, POA, \text{ or } SOA$ ) and fitting parameters are shown as multivariate  
554 probability density functions in Fig. S9. This yielded  $\alpha_{BC} = 1.2$ ,  $\alpha_{POA} = 4.6$ , and  $\alpha_{SOA} = 5.6$ , with corresponding  
555 uncertainties of approximately 20% (complete details of the uncertainties are provided in Table S1). Note that  $\alpha_{BC}$  in  
556 the range 660 nm to 370 nm obtained from this fit is very similar to  $\alpha_{BC}$  values that can be inferred by extrapolating  
557 the data shown in Fig. 3A to  $f_{OA}=0$ . ~~The high  $\alpha$  values obtained for the organic fractions are consistent with previous~~  
558 ~~measurements for BrC-containing POA (e.g. Chakrabarty et al., 2010, 2013) although, to our knowledge, this is the~~  
559 ~~first study to report  $\alpha_{SOA}$  without performing a solvent extraction.~~ The high  $\alpha$  values obtained for the organic  
560 fractions are consistent with previous measurements for BrC containing POA (e.g. Chakrabarty et al., 2010, 2013).

561 **Evolution of  $MAC_{OA}$  with aging.** In Fig. 5, we examine whether the absorption profile of SOA evolved with aging.  
562 A change in  $MAC_{SOA}(370\text{nm})$  or  $\alpha_{SOA}$  with increasing OH exposure may indicate either a change in the mass-  
563 specific absorption of the condensing SOA species with time, or a change (e.g. “bleaching”) in the MAC of pre-  
564 existing POA. Fig. 5 indicates that neither of these scenarios was the case. Both  $MAC_{SOA}(370\text{nm})$  and  $\alpha_{SOA}$  were  
565 statistically independent of the OH exposure, for exposures up to 40 molec. OH  $\text{cm}^{-3}$  h. This signifies that under our  
566 conditions and within our measurement uncertainties the optical properties of the additional organic mass formed  
567 was constant with aging, under the assumption that the light-absorption properties of POA were negligibly

568 influenced by aging. Most of the variability in  $MAC_{SOA}(\lambda)$  discussed above is therefore related to experiment-to-  
569 experiment differences rather than to the extent of OH exposure, as it is also shown below.

### 570 4.3 Solubility of BrC in methanol and water

571 Fig. 6 shows the  $MAC_{OA}(370nm)$  determined from the water and methanol extracts against the  $MAC_{OA}(370nm)$   
572 determined from the online measurements. The  $MAC_{OA}(370nm)$  from online measurements was estimated by  
573 subtracting the contribution of BC assuming a constant  $MAC_{BC}(370nm) = 13.7 \text{ m}^2 \cdot \text{g}^{-1}$  as obtained in this work  
574 (Table 1). We performed all the calculations and comparisons at  $\lambda = 370 \text{ nm}$ , as the signal to noise ratio of the  
575 absorption coefficients measured by UV-visible spectroscopy and the contribution of BrC to the total carbonaceous  
576 absorption are highest at this wavelength. The MAC of the extracts was computed from the  $k_{OA}$  through Mie  
577 calculations. Repetition of both water and methanol extracts yielded results that were consistent within 10% (Fig.  
578 S11). Average raw absorption spectra are shown in Fig. S12.

579 ~~Fig. 6B shows excellent correlation between the MAC values obtained from the the  $k_{OA}$  of the methanol extracted~~  
580 ~~OA with the in situ method described above.~~ Fig. 6B shows excellent correlation between the  $MAC_{OA}(370nm)$   
581 values obtained from the  $k_{OA}$  of the solvent-extracted OA with the in-situ method described above. The Pearson  
582 correlation coefficient was 0.8, for both solvents. This correlation suggests that none of the assumptions employed in  
583 either method led to substantial errors in precision, providing direct support for our results. A similar relationship  
584 was observed between  $k_{OA}$  and the  $MAC_{OA}(370nm)$  determined from the online measurements (Fig. S13), showing  
585 that this relationship is not sensitive to assumptions underlying the Mie calculations. It further suggests that the  
586 wide variability observed in the  $MAC_{OA}$  values of different burns, as seen Fig. 6, most likely reflects real variability  
587 in the optical properties of POA and SOA rather than random noise or experimental errors in the retrieved quantities.  
588  $MAC_{OA}$  retrieved based on the  $k_{OA}$  of the water soluble OA show substantially more scatter than observed in Fig. 6B  
589 (for both primary and aged data), suggesting a variable extraction efficiency in the case of water, which we also  
590 attribute to variability in the OA composition.

591 ~~The orthogonal, uncertainty weighted linear regression in Fig. 6B shows that the methanol extracts explain 46~~  
592  ~~$\pm 10\%$  of the online MAC. (Note that, in this analysis, aged OA refers to the sum of POA and SOA for aged~~  
593 ~~samples.) Considering the simplifying assumptions that were necessary for our Mie analysis and those related to~~  
594 ~~online  $MAC_{OA}$  calculations, we consider this an adequate agreement. In particular, the assumption of a perfect~~  
595 ~~extraction efficiency of OA in methanol may have been violated (see Section 3.4). Conversely, the fit in Fig. 6A~~

596 indicates that the apparent MAC of water soluble species was a fourth of the respective methanol MAC, according  
597 to the slope of only  $12 \pm 3\%$ . This strong disagreement shows that the BrC in our samples was hardly water  
598 soluble, even for the most aged samples. As we expect that the majority of OA in our samples formed by wood  
599 pyrolysis (Di Blasi, 2008; Corbin et al., 2015b; Shafizadeh, 1984), we can compare our results directly to those of  
600 Chen and Bond (2010), who also found that primary wood pyrolysis BrC was water insoluble. Moreover, the water-  
601 insoluble nature of the light absorbing components of SOA is in line with the results by Bruns et al. (2016) who  
602 showed that the precursors of SOA in these experiments were predominantly aromatic compounds.

603 The data in Fig. 6B show that the methanol extracts correspond to a MAC about 50% smaller than the online data.  
604 The scatter in the data is significantly reduced for the aged data (note that, in this analysis, aged OA refers to the  
605 sum of POA and SOA, since the reported values represent all OA after aging). This reduced scatter is expected,  
606 considering that aging is likely to result in more-spherical particles. We have assumed particle sphericity when  
607 interpreting the SMPS data and performing the Mie analysis. While the propagation of quantifiable uncertainties  
608 leads to an error estimate of  $\sim 25\%$ , considering the simplifications that were necessary for the Mie analysis, we  
609 consider a 50% closure to be an adequate agreement. Despite this, we cannot exclude additional methanol insoluble  
610 brown carbon. Conversely, the fit in Fig. 6A indicates that the apparent MAC of water-soluble species was a fourth  
611 of the respective methanol MAC, according to the slope of only  $12 \pm 3\%$ . Only the aged data have been fit to  
612 illustrate this point. This strong disagreement shows that the BrC in our samples was hardly water soluble, even for  
613 the most aged samples. As we expect that the majority of OA in our samples formed by wood pyrolysis (Di Blasi,  
614 2008; Corbin et al., 2015b; Shafizadeh, 1984), we can compare our results directly to those of Chen and Bond  
615 (2010), who also found that primary wood-pyrolysis BrC was water insoluble. Moreover, the poor water solubility  
616 of the light absorbing components of SOA (Zhang et al., 2011) is in line with the results by Bruns et al. (2016) who  
617 showed that SOA precursors during these experiments were predominantly aromatic compounds.

#### 618 4.4 Comparison of $k_{OA}$ with literature

619 The results above highlight the variability in the OA absorption properties. In this section, we discuss potential  
620 reasons for this variability and compare our results to literature. Fig. 7 shows the imaginary refractive index of  
621 methanol-extracted OA at 370 nm,  $k_{OA, \text{methanol}}(370\text{nm})$  (Eq. (8)), as a function of  $M_{BC}/M_{OA}$  and aging. The data are  
622 plotted against  $M_{BC}/M_{OA}$  instead of  $f_{OA}$  to allow for a direct comparison with literature (see Fig. S14 for a plot  
623 against  $f_{OA}$ ). An approximately linear trend of  $k_{OA, \text{methanol}}(370\text{nm})$  with  $M_{BC}/M_{OA}$  is seen in log space. This aging-

624 independent relationship may be useful in, for example, atmospheric scenarios where wood-burning OA is a  
625 dominant aerosol component but its exact degree of aging is unknown. The decrease of  $M_{BC}/M_{OA}$  caused by  
626 formation of SOA during aging results in a concurrent decrease of  $k_{OA, \text{methanol}}(370\text{nm})$ , implying that  $k_{SOA} < k_{POA}$ .  
627 This result is consistent with the smaller MAC of SOA compared to POA obtained from online measurements  
628 (Table 1) and with recent results reported by Sumlin et al. (2017). We emphasize that the derived quantity here is  
629 the imaginary refractive index  $k$  of the total aged OA, not the SOA.

630 The increase of  $k_{OA, \text{methanol}}(370\text{nm})$  with increasing  $M_{BC}/M_{OA}$  indicates that the OA compounds present at higher  
631  $M_{BC}/M_{OA}$  absorbed more efficiently than at low  $M_{BC}/M_{OA}$ . If the variability in  $M_{BC}/M_{OA}$  was driven partly by OA  
632 partitioning, then this implies that lower-volatility compounds were more absorbing than high-volatility compounds,  
633 consistent with the results by Saleh et al. (2014) who investigated the relation between OA absorption and volatility  
634 using thermodesorber measurements. A correlation between  $k_{OA}$  and  $M_{BC}/M_{OA}$  has also been reported by Lu et al.  
635 (2015). ~~The parameterizations reported by these authors are included in Fig. 7, both showing a smaller trend with~~  
636  ~~$M_{BC}/M_{OA}$  than seen in our data.~~ The parameterizations reported by these authors are included in Fig. 7, where the  
637 wavelength dependence reported by those authors has been used to adjust their parameterizations to 370nm. Despite  
638 these differences, our results confirm the generality of the correlation proposed by Saleh et al. (2014), but using a  
639 method that is independent of potential biases related to internal mixing effects, filter-based absorption  
640 measurements or Mie calculations. Indeed, we emphasize that the  $k_{OA}$  obtained here is a lower limit: as our approach  
641 does not account for the OA extraction efficiency;  $k_{OA, \text{methanol}}(370\text{nm})$  may be underestimated by up to a factor of ~2,  
642 based on Fig. 6B.

643

## 644 5. ATMOSPHERIC IMPLICATIONS

645 In this section, we seek to estimate the relative importance of OA absorption at different wavelengths relative to that  
646 of the total carbonaceous aerosol as a function of aging. For these calculations, the  $MAC(\lambda)$  values for the different  
647 components and their relative mass abundance are required. We used the power law parameters reported above to  
648 generate continuous  $MAC_{BC}(\lambda)$ ,  $MAC_{POA}(\lambda)$ , and  $MAC_{SOA}(\lambda)$  functions together with their associated uncertainties  
649 (Fig. 8A), which allow the extrapolation of these parameters in the range [280nm; 880nm].

650 The contributions of the different components as a function of OH exposure were calculated by assuming that SOA  
651 production follows the first order decay of its precursors, i.e., the reaction with OH. Under this assumption, the time-  
652 dependent mass concentration of SOA compared to POA can be expressed as

$$653 \quad M_{\text{SOA,WLC}}(t)/M_{\text{POA,WLC}}(t) = M_{\text{SOAP,WLC}}/M_{\text{POA,WLC}} \times (1 - \exp(-k_{\text{OH}}OH_{\text{exp}})) \quad (21)$$

654 In this equation,  $M_{\text{SOA,WLC}}(t)$ ,  $M_{\text{POA,WLC}}(t)$  and  $M_{\text{SOAP,WLC}}$  are the wall loss corrected mass concentrations of SOA,  
655 POA and the SOA potential (the maximum SOA formed upon the consumption of all precursors).  $k_{\text{OH}}$  represents an  
656 estimation of reaction rate of SOA precursors towards OH based on SOA production rates. By fitting the observed  
657  $M_{\text{SOA,WLC}}(t)/M_{\text{POA,WLC}}(t)$  against the OH exposure,  $k_{\text{OH}}$  and  $M_{\text{SOAP,WLC}}/M_{\text{POA,WLC}}$  can be estimated. For these  
658 calculations, we have estimated the wall losses using two approaches as described in the SI.

659 ~~The  $M_{\text{SOAP,WLC}}/M_{\text{POA,WLC}}$  was on average equal to 7.8 (GSD = 1.4) and  $k_{\text{OH}}$  was estimated as  $2.7 \times 10^{-11}$  molecule<sup>-1</sup>~~  
660 ~~cm<sup>3</sup> (GSD = 1.4), consistent with the chemically speciated data obtained by a proton-transfer-reaction mass~~  
661 ~~spectrometer (PTR-MS) (Bruns et al., 2016, 2017). These high rates and enhancement ratios indicate the rapid~~  
662 ~~production of SOA.~~ The  $M_{\text{SOAP,WLC}}/M_{\text{POA,WLC}}$  was on average equal to 7.8 (GSD = 1.4) and  $k_{\text{OH}}$  was estimated as  
663  $2.7 \times 10^{-11}$  molecule<sup>-1</sup> cm<sup>3</sup> (GSD = 1.4), consistent with the SOA precursors chemical nature measured (e.g. PAH and  
664 phenol derivatives) by a proton-transfer-reaction mass spectrometer (PTR-MS) (Bruns et al., 2016, 2017). These  
665 high rates and enhancement ratios indicate the rapid production of SOA.

666 Based on the bulk gas phase measurements of SOA precursors (Bruns et al., 2016), the obtained enhancements are  
667 consistent with high bulk SOA yields of ~50%. These high values are not surprising, considering the nature of these  
668 gases (e.g. PAH and phenol derivatives), the low temperatures (263 K), and the relatively high concentrations (Aged  
669 OA ~100 µg m<sup>-3</sup>) at which the experiments have been conducted (Bruns et al. 2016).

670 Combining these calculated enhancements with the average contributions of POA in primary emissions, the  
671 evolution of  $f_{\text{OA}}$  with aging was determined and is shown in Fig. 8B. The uncertainties in Fig. 8B (dotted lines)  
672 represent one standard deviation on  $f_{\text{OA}}$  obtained by a Monte Carlo propagation of uncertainties due to experiment-  
673 to-experiment variability, fitting errors and wall loss correction errors (see SI). While this calculation represents a  
674 simplification of the SOA production mechanisms (the dependence of SOA yields on OH exposures/multigeneration  
675 chemistry and OA mass concentrations was neglected), it results in residuals much smaller than the experiment-to-  
676 experiment variability. We therefore used these calculations to assess the relative contribution of OA to the total

677 carbonaceous absorption. We show in Fig. 8C that below 400 nm and upon aging, the absorption coefficient of the  
678 total organics was at least as high as the one of BC.

679 Using the MAC values of the different components (in  $\text{m}^2 \text{g}^{-1}$ ), their abundance (in  $\text{g m}^{-3}$ ) and the solar irradiance  
680 data ( $S$ , in  $\text{W m}^{-2} \text{nm}^{-1}$ ) calculated at sea level for a cloudless day, the fractional energy transfer due to the BrC light  
681 absorption relative to that due to the total carbonaceous aerosol absorption,  $W_{\text{OA}}(\text{OH}_{\text{exp}})$ , in air masses dominated  
682 by residential burning emissions can be determined as

$$\begin{aligned} 683 \quad W_{\text{OA}}(\text{OH}_{\text{exp}}) &= \text{RET}_{\text{OA}}(\text{OH}_{\text{exp}}) / \text{RET}_{\text{tot}}(\text{OH}_{\text{exp}}) \\ 684 \quad &= \frac{\int_{300}^{880} \{M_{\text{POA}}(\text{OH}_{\text{exp}}) \times \text{MAC}_{\text{POA}}(\lambda) + M_{\text{SOA}}(\text{OH}_{\text{exp}}) \times \text{MAC}_{\text{SOA}}(\lambda)\} \times S(\lambda) \times d\lambda}{\int_{300}^{880} \{M_{\text{BC}}(\text{OH}_{\text{exp}}) \times \text{MAC}_{\text{BC}}(\lambda) + M_{\text{POA}}(\text{OH}_{\text{exp}}) \times \text{MAC}_{\text{POA}}(\lambda) + M_{\text{SOA}}(\text{OH}_{\text{exp}}) \times \text{MAC}_{\text{SOA}}(\lambda)\} \times S(\lambda) \times d\lambda} \quad (22) \end{aligned}$$

685 Here,  $\text{RET}_{\text{OA}}(\text{OH}_{\text{exp}})$  and  $\text{RET}_{\text{tot}}(\text{OH}_{\text{exp}})$  denote the rate of energy transfer per volume (in  $\text{W m}^{-3}$ ) to the air mass in  
686 question due to light absorption by OA and the total carbonaceous aerosol, respectively. We note that while  
687  $\text{RET}_{\text{OA}}(\text{OH}_{\text{exp}})$  and  $\text{RET}_{\text{tot}}(\text{OH}_{\text{exp}})$  are extensive properties,  $W_{\text{OA}}(\text{OH}_{\text{exp}})$  does not depend on the loading or  
688 scattering/lensing, provided that scattering/lensing similarly affects BC and OA present in the same air mass (e.g.  
689 BC and OA have a similar size distribution).

690 We also note that  $W_{\text{OA}}(\text{OH}_{\text{exp}})$  depends on the photon flux,  $S(\lambda)$ , but we consider this dependence to be trivial  
691 compared to the variability in the aerosol emissions and their light absorbing properties (error bars considering these  
692 variabilities are shown in Fig. 8D). Errors in  $W_{\text{OA}}$  were propagated by Monte Carlo simulations using the  
693 uncertainties from the estimated MAC values of BC and OA fractions and the variability in  $f_{\text{OA}}$ . Our ~~error~~  
694 **sensitivity** analysis suggests that the major part of the variance in predicting  $W_{\text{OA}}$  for primary emissions stems from  
695 the variability in the POA mass fraction. In contrast, the SOA mass absorption cross-sections at lower wavelengths  
696 are the most critical factor for assessing the relative importance of BrC absorptivity in aged emissions.

697 Fig. 8D shows that the fractional energy transfer to the air mass,  $W_{\text{OA}}$ , due to the absorption by the primary organic  
698 aerosol was around 10% of that of the total carbonaceous aerosol for our samples. This percentage is comparable to  
699 that observed by Fu et al. (2012), in spite of  $f_{\text{OA}}$  in their samples being much higher, because of the high OA MACs  
700 in our samples (Table 1). Moreover, with aging, the fraction of OA is enhanced, resulting in a sizeable increase  $W_{\text{OA}}$ ,  
701 from ~0.1 to ~0.3 (Fig. 8D), highlighting that SOA formation in biomass burning plumes is an atmospherically  
702 relevant source of BrC. We note that our data are more representative of flaming conditions. More data are needed



703 on the chemical nature of primary particulate emissions and of the contributing SOA precursors as well as the  
704 absorptivity of these primary and secondary products, for better constraining the influence of biomass-burning  
705 related BrC on the Earth's climate.

706

## 707 6. CONCLUSIONS

708 We determined wavelength-dependent MAC values of BC, POA and SOA, as well as  $k_{\text{OA}}$  for methanol and water  
709 extracts of fresh and aged OA, for wood-burning emissions through smog-chamber experiments. To our knowledge,  
710 this is the first determination of these properties for wood-burning OA. We showed that the  $\text{MAC}_{\text{OA}}(370\text{nm})$  values  
711 calculated based on  $k_{\text{OA}}$  through Mie analysis correlated well with those estimated from online filter based  
712 measurements. This correlation between independent MAC measurements supports the quality of both methods.  
713 While  $\text{MAC}_{\text{OA}}(370\text{nm})$  values computed based on  $k_{\text{OA, methanol}}$  were a 2-fold lower than those estimated from online  
714 filter based measurements, calculations based on  $k_{\text{OA, water}}$  could only explain 12% of the measured absorption,  
715 suggesting that BrC species in POA and SOA are mostly water insoluble. The  $\text{MAC}_{\text{OA}}$  was found to vary by more  
716 than one order of magnitude. Similar to previous reports, this variability could be related to the variability in the  
717 ratio of the mass concentrations of BC and OA ( $M_{\text{BC}}/M_{\text{OA}}$ ) due to very different mechanisms of oxidative aging and  
718 burn-to-burn variability.

719 The  $\text{MAC}_{\text{POA}}$  and  $\text{MAC}_{\text{SOA}}$  determined for wavelengths between 370 and 660 nm followed a power-law dependence  
720 on  $\lambda$  with an absorption Ångström exponent of 4.6 and 5.6 for POA and SOA, respectively. In addition to following  
721 this power law, the MACs of POA and SOA appeared to be constant for OH exposures up to  $40 \times 10^6$  molecules  $\text{cm}^{-3}$   
722 h.

723 The mean  $\text{MAC}_{\text{POA}}(370\text{nm})$  obtained under our conditions was  $5.5 \text{ m}^2 \text{ g}^{-1}$ , considerably higher than previously  
724 reported values for open biomass burning. ~~The mean  $\text{MAC}_{\text{SOA}}(370\text{nm})$  was  $2.4 \text{ m}^2 \text{ g}^{-1}$  under our experimental~~  
725 ~~conditions, 2.3 times lower than the mean  $\text{MAC}_{\text{POA}}(370\text{nm})$  but approximately an order of magnitude higher than~~  
726 ~~MAC values estimated for ambient oxygenated aerosols or reported for SOA from biogenic and traditional~~  
727 ~~anthropogenic precursors. The mean  $\text{MAC}_{\text{SOA}}(370\text{nm})$  was  $2.2 \text{ m}^2 \text{ g}^{-1}$  (one-sigma variability:  $1.6 - 3.1 \text{ m}^2 \text{ g}^{-1}$~~   
728 ~~according to a GSD = 1.39) under our experimental conditions, 2.3 times lower than the mean  $\text{MAC}_{\text{POA}}(370\text{nm})$  but~~  
729 ~~approximately an order of magnitude higher than MAC values estimated for ambient oxygenated aerosols or~~

730 reported for SOA from biogenic and traditional anthropogenic precursors. We propose that the important role of  
731 oxidized phenols and aromatics in forming wood-burning SOA (Brunns et al., 2016) is the cause of this observation.  
732 This hypothesis is supported by our observed reaction rates with OH, and by the water-insolubility of the BrC in  
733 aged OA.

734 Overall, the absorption by organic aerosols was estimated to contribute 10-30% of the total solar absorption of  
735 wood-combustion aerosols, where 10% represents the primary OA and 30% the aged OA. SOA formation in  
736 biomass burning plumes is therefore an atmospherically relevant source of BrC.

737 *Acknowledgements.* The research leading to these results has received funding from the European Research Council  
738 grant (ERC-CoG 615922-BLACARAT) and by the Competence Centre Energy and Mobility (CCEM) project 807.

739

740

741

742

743

744

745

746

747

748

749

750

751

752

753

754

755

756

757

758

759 **References**

760 Alexander, D. T. L., Crozier, P. A. and Anderson, J. R.: Brown carbon spheres in East Asian outflow and their  
761 optical properties., *Science*, 321(5890), 833–6, doi:10.1126/science.1155296, 2008.

762 Andreae, M. O. and Gelencsér, A.: Black carbon or brown carbon? the nature of light-absorbing carbonaceous  
763 aerosols, *Atmos. Chem. Phys.*, 6(3), 3419–3463, doi:10.5194/acpd-6-3419-2006, 2006.

764 Barmet, P., Dommen, J., DeCarlo, P. F., Tritscher, T., Praplan, A. P., Platt, S. M., Prévôt, A. S. H., Donahue, N. M.  
765 and Baltensperger, U.: OH clock determination by proton transfer reaction mass spectrometry at an environmental  
766 chamber, *Atmos. Meas. Tech.*, 5(3), 647–656, doi:10.5194/amt-5-647-2012, 2012.

767 Bertrand, A., Stefanelli, G., Bruns, E. A., Pieber, S. M., Prévôt, A. S. H., Wortham, H., Temime-roussel, B.,  
768 Slowik, J. G., Haddad, I. El and Marchand, N.: Primary emissions and secondary aerosol production potential from  
769 woodstoves for residential heating : influence of the stove technology and combustion efficiency, *Atmos. Environ.*,  
770 169, doi:10.1016/j.atmosenv.2017.09.005, 2017.

771 Di Blasi, C.: Modeling chemical and physical processes of wood and biomass pyrolysis, *Prog. Energy Combust.*  
772 *Sci.*, 34(1), 47–90, doi:10.1016/j.pecs.2006.12.001, 2008.

773 Bond, T. C., Habib, G. and Bergstrom, R. W.: Limitations in the enhancement of visible light absorption due to  
774 mixing state, *J. Geophys. Res. Atmos.*, 111(20), 1–13, doi:10.1029/2006JD007315, 2006.

775 Bond, T. C., Doherty, S. J., Fahey, D. W., Forster, P. M., Berntsen, T., Deangelo, B. J., Flanner, M. G., Ghan, S.,  
776 Kärcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P. K., Sarofim, M. C., Schultz, M. G., Schulz, M.,  
777 Venkataraman, C., Zhang, H., Zhang, S., Bellouin, N., Guttikunda, S. K., Hopke, P. K., Jacobson, M. Z., Kaiser, J.  
778 W., Klimont, Z., Lohmann, U., Schwarz, J. P., Shindell, D., Storelvmo, T., Warren, S. G. and Zender, C. S.:  
779 Bounding the role of black carbon in the climate system: a scientific assessment, *J. Geophys. Res. Atmos.*, 118(11),  
780 5380–5552, doi:10.1002/jgrd.50171, 2013.

781 Bruns, E. A., Krapf, M., Orasche, J., Huang, Y., Zimmermann, R., Drinovec, L., Močnik, G., El-Haddad, I., Slowik,  
782 J. G., Dommen, J., Baltensperger, U. and Prévôt, A. S. H.: Characterization of primary and secondary wood  
783 combustion products generated under different burner loads, *Atmos. Chem. Phys.*, 15(5), 2825–2841,  
784 doi:10.5194/acp-15-2825-2015, 2015.

785 Bruns, E. A., El Haddad, I., Slowik, J. G., Kilic, D., Klein, F., Baltensperger, U. and Prévôt, A. S. H.: Identification  
786 of significant precursor gases of secondary organic aerosols from residential wood combustion., *Sci. Rep.*, 6, 27881,  
787 doi:10.1038/srep27881, 2016.

788 Bruns, E. A., Slowik, J. G., Haddad, I. El, Kilic, D., Klein, F., Dommen, J., Temime-Roussel, B., Marchand, N.,  
789 Baltensperger, U. and Prévôt, A. S. H.: Characterization of gas-phase organics using proton transfer reaction time-  
790 of-flight mass spectrometry : fresh and aged residential wood combustion emissions, *Atmos. Chem. Phys.*, 705–720,  
791 doi:10.5194/acp-17-705-2017, 2017.

792 Cavalli, F., Viana, M., Yttri, K. E., Genberg, J. and Putaud, J.-P.: Toward a standardised thermal-optical protocol for  
793 measuring atmospheric organic and elemental carbon: the EUSAAR protocol, *Atmos. Meas. Tech.*, 3(1), 79–89,  
794 doi:doi:10.5194/amt-3-79-2010, 2010.

795 Chakrabarty, R. K., Moosmüller, H., Chen, L. W. A., Lewis, K., Arnott, W. P., Mazzoleni, C., Dubey, M. K., Wold,  
796 C. E., Hao, W. M. and Kreidenweis, S. M.: Brown carbon in tar balls from smoldering biomass combustion, *Atmos.*  
797 *Chem. Phys.*, 10(13), 6363–6370, doi:10.5194/acp-10-6363-2010, 2010.

798 Chakrabarty, R. K., Arnold, I. J., Francisco, D. M., Hatchett, B., Hosseinpour, F., Loria, M., Pokharel, A. and  
799 Woody, B. M.: Black and brown carbon fractal aggregates from combustion of two fuels widely used in asian  
800 rituals, *J. Quant. Spectrosc. Radiat. Transf.*, 122, 25–30, doi:10.1016/j.jqsrt.2012.12.011, 2013.

801 Chen, Y. and Bond, T. C.: Light absorption by organic carbon from wood combustion, *Atmos. Chem. Phys.*, 10,  
802 1773-1787, doi:10.5194/acp-10-1773-2010, 2010.

803 China, S., Mazzoleni, C., Gorkowski, K., Aiken, A. C., and Dubey, M. K.: Morphology and mixing state of  
804 individual freshly emitted wildfire carbonaceous particles, *Nat. Commun.*, 4, 1–7, doi:10.1038/ncomms-3122-2013,  
805 2013.

806 Ciarelli, G., Haddad, I. El, Bruns, E. and Aksoyoglu, S., Möhler, O., Baltensperger, U. and Prévôt, A. S. H.:  
807 Constraining a hybrid volatility basis set model for aging of wood burning emissions using smog chamber  
808 experiments, *Geosci. Model. Dev.*, 2303-2320, doi:10.5194/gmd-2303-2017, 2017.

809 Clarke, A., McNaughton, C., Kapustin, V., Shinozuka, Y., Howell, S., Dibb, J., Zhou, J., Anderson, B. E.,  
810 Brekhovskikh, V., Turner, H. and Pinkerton, M.: Biomass burning and pollution aerosol over North America:  
811 organic components and their influence on spectral optical properties and humidification response, *J. Geophys. Res.*  
812 *Atmos.*, 112(12), 1–13, doi:10.1029/2006JD007777, 2007.

813 Collaud Coen, M., Weingartner, E., Apituley, A., Ceburnis, D., Fierz-Schmidhauser, R., Flentje, H., Henzing, J. S.,  
814 Jennings, S. G., Moerman, M., Petzold, A., Schmid, O., and Baltensperger, U.: Minimizing light absorption  
815 measurement artifacts of the Aethalometer: evaluation of five correction algorithms, *Atmos. Meas. Tech.*, 3, 457–  
816 474, doi:10.5194/amt-3-457-2010, 2010.

817  
818 Corbin, J. C., Lohmann, U., Sierau, B., Keller, A., Burtscher, H. and Mensah, A. A.: Black carbon surface oxidation  
819 and organic composition of beech-wood soot aerosols, *Atmos. Chem. Phys.*, 15(20), 11885–11907, doi:10.5194/acp-  
820 15-11885-2015, 2015a.

821 Corbin, J. C., Keller, A., Lohmann, U., Burtscher, H., Sierau, B. and Mensah, A. A.: Organic emissions from a wood  
822 stove and a pellet stove before and after simulated atmospheric aging, *Aerosol Sci. Technol.*, 49(11), 1037–1050,  
823 doi:10.1080/02786826.2015.1079586, 2015b.

824 Corbin, J. C.; Pieber, S. M.; Czech, H.; Zanatta, M.; Jakobi, G.; Massabò, D.; Orasche, J.; El Haddad, I.; Mensah, A.  
825 A.; Stengel, B.; Drinovec, L.; Mocnik, G.; Zimmermann, R.; Prévôt, A. S. H. and Gysel, M.: Brown and black  
826 carbon emitted by a marine engine operated on heavy fuel oil and distillate fuels: optical properties, size  
827 distributions and emission factors, *J. Geophys. Res. Atmos.*, 123, 6175-6195, doi:10.1029/2017JD027818, 2018.

828 DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C., Gonin, M., Fuhrer, K.,  
829 Horvath, T., Docherty, K. S., Worsnop, D. R. and Jimenez, J. L.: Field deployable, high resolution, time-of-flight  
830 aerosol mass spectrometer, *Anal. Chem.*, 78(24), 8281–8289, doi:10.1029/2001JD001213, 2006.

831 Denier Van Der Gon, H. A. C., Bergström, R., Fountoukis, C., Johansson, C., Pandis, S. N., Simpson, D. and  
832 Visschedijk, A. J. H.: Particulate emissions from residential wood combustion in Europe - revised estimates and an  
833 evaluation, *Atmos. Chem. Phys.*, 15(11), 6503–6519, doi:10.5194/acp-15-6503-2015, 2015.

834 Drinovec, L., Močnik, G., Zotter, P., Prévôt, A. S. H., Ruckstuhl, C., Coz, E., Rupakheti, M., Sciare, J., Müller, T.,  
835 Wiedensohler, A. and Hansen, A. D. A.: The “dual-spot” aethalometer: an improved measurement of aerosol black  
836 carbon with real-time loading compensation, *Atmos. Meas. Tech.*, 8(5), 1965–1979, doi:10.5194/amt-8-1965-2015,  
837 2015.

838 Feng, Y., Ramanathan, V. and Kotamarthi, V. R.: Brown carbon: a significant atmospheric absorber of solar  
839 radiation, *Atmos. Chem. Phys.*, 13(17), 8607–8621, doi:10.5194/acp-13-8607-2013, 2013.

840 Forrister, H., Liu, J., Scheuer, E., Dibb, J., Ziemba, L., Thornhill, L. K., Anderson, B., Diskin, G., Perring, A. E.,  
841 Schwarz, J. P., Campuzan-Jost, P., Day, D. A., Palm, B. B., Jimenez, J. L., Nenes, A., Weber, R. J.: Evolution of  
842 brown carbon in wildfire plumes, *Gephys. Res. Lett.*, 42, 4623-4630, doi: 10.1002/2015GL063897, 2015.

843  
844 Fu, J. S., Hsu, N. C., Gao, Y., Huang, K., Li, C., Lin, N. H. and Tsay, S. C.: Evaluating the influences of biomass  
845 burning during 2006 BASE-ASIA: a regional chemical transport modeling, *Atmos. Chem. Phys.*, 12(9), 3837–3855,  
846 doi:10.5194/acp-12-3837-2012, 2012.

- 847 Grieshop, A. P., Logue, J. M., Donahue, N. M. and Robinson, A. L.: Laboratory investigation of photochemical  
848 oxidation of organic aerosol from wood fires – part 1: measurement and simulation of organic aerosol evolution,  
849 *Atmos. Chem. Phys.*, 9, 2227–2240, doi:10.5194/acp-9-2227-2009, 2009.
- 850 Gueymard, C.; Myers, D.; Emery, K.: Proposed Reference Irradiance Spectra for Solar Energy Systems Testing,  
851 *Solar Energy*, 73, 6, 443–467, 2002.  
852
- 853 Gundel, L. A., Dod, R. L., Rosen, H. and Novakov.: The relationship between optical attenuation and black carbon  
854 concentration for ambient and source particles, *Sci. Total Environ.*, 36, 197-202, 1984.
- 855 Haslett, S. L.; Thomas, J. C.; Morgan, W. T.; Hadden, R.; Liu, D.; Allan, J. D.; Williams, P. I.; Keita, S.; Lioussé, C.  
856 and Coe, H.: Highly controlled, reproducible measurements of aerosol emissions from combustion of a common  
857 African biofuel source, *Atmos. Chem. Phys.*, 385-403, doi:10.5194/acp-18-385-2018, 2018.  
858
- 859 Hecobian, A., Zhang, X., Zheng, M., Frank, N., Edgerton, E. S. and Weber, R. J.: Water-soluble organic aerosol  
860 material and the light-absorption characteristics of aqueous extracts measured over the southeastern United States,  
861 *Atmos. Chem. Phys.*, 10(13), 5965–5977, doi:10.5194/acp-10-5965-2010, 2010.
- 862 Heringa, M. F., DeCarlo, P. F., Chirico, R., Tritscher, T., Dommen, J., Weingartner, E., Richter, R., Wehrle, G.,  
863 Prévôt, A. S. H. and Baltensperger, U.: Investigations of primary and secondary particulate matter of different wood  
864 combustion appliances with a high-resolution time-of-flight aerosol mass spectrometer, *Atmos. Chem. Phys.*,  
865 11(12), 5945–5957, doi:10.5194/acp-11-5945-2011, 2011.
- 866 Hoffer, A., Gelencsér, A., Guyon, P., Kiss, G., Schmid, O., Frank, G. P., Artaxo, P. and Andreae, M. O.: Optical  
867 properties of humic-like substances (HULIS) in biomass-burning aerosols, *Atmos. Chem. Phys.*, 6, 3563-3570,  
868 doi:10.5194/acp-6-3563-2006, 2006.
- 869 Jo, D. S., Park, R. J., Lee, S., Kim, S. W. and Zhang, X.: A global simulation of brown carbon: Implications for  
870 photochemistry and direct radiative effect, *Atmos. Chem. Phys.*, 16(5), 3413–3432, doi:10.5194/acp-16-3413-2016,  
871 2016.
- 872 Kirchstetter, T. W., Novakov, T. and Hobbs, P. V.: Evidence that the spectral dependence of light absorption by  
873 aerosols is affected by organic carbon, *J. Geophys. Res. D Atmos.*, 109(21), 1–12, doi:10.1029/2004JD004999,  
874 2004.
- 875 Krapf, M., Haddad, I. El, Bruns, E. A., Haddad, I. El, Molteni, U., Daellenbach, K. R., Prévôt, A. S. H.,  
876 Baltensperger, U., Dommen, J. : Labile peroxides in secondary organic aerosol labile peroxides in secondary organic  
877 aerosol, *Chem 1*, 603–616, doi:10.1016/j.chempr.2016.09.007, 2016.
- 878 Lack, D. A. and Langridge, J. M.: On the attribution of black and brown carbon light absorption using the Ångström  
879 exponent, *Atmos. Chem. Phys.*, 13(20), 10535–10543, doi:10.5194/acp-13-10535-2013, 2013.
- 880 Lambe, A. T., Cappa, C. D., Massoli, P., Onasch, T. B., Forestieri, S. D., Martin, A. T., Cummings, M. J.,  
881 Croasdale, D. R., Brune, W. H., Worsnop, D. R. and Davidovits, P.: Relationship between oxidation level and  
882 optical properties of secondary organic aerosol, *Environ. Sci. Technol.*, 47(12), 6349–6357, doi:10.1021/es401043j,  
883 2013.
- 884 Laskin, A., Laskin, J. and Nizkorodov, S. A.: Chemistry of atmospheric brown carbon, *Chem. Rev.*, 115(10), 4335–  
885 4382, doi:10.1021/cr5006167, 2015.
- 886 Lee, H. J., Aiona, P. K., Laskin, A., Laskin, J. and Nizkorodov, S. A.: Effect of solar radiation on the optical  
887 properties and molecular composition of laboratory proxies of atmospheric brown carbon, *Environ. Sci. Technol.*,  
888 48(17), 10217–10226, doi:10.1021/es502515r, 2014.
- 889 Lewis, K., Arnott, W. P., Moosmüller, H. and Wold, C. E.: Strong spectral variation of biomass smoke light  
890 absorption and single scattering albedo observed with a novel dual-wavelength photoacoustic instrument, *J.*

- 891 Geophys. Res. Atmos., 113(16), 1–14, doi:10.1029/2007JD009699, 2008.
- 892 Lin, G., Penner, J. E., Flanner, M. G., Sillman, S., Xu, L. and Zhou, C.: Radiative forcing of organic aerosol in the  
893 atmosphere and on snow: effect of SOA and brown carbon, *J. Geophys. Res. - Atmos.*, 119(12), 7453–7476,  
894 doi:10.1002/2013JD021186. Received, 2014.
- 895 Liu, J., Scheuer, E., Dibb, J., Ziemba, L. D., Thornhill, K. L., Anderson, B. E., Wisthaler, A., Mikoviny, T., Devi, J.  
896 J., Bergin, M. and Weber, R. J.: Brown carbon in the continental troposphere, *Geophys. Res. Lett.*, 41, 2191–2195,  
897 doi:10.1002/2013GL058976, 2014.
- 898 Liu, J., Lin, P., Laskin, A., Laskin, J., Kathmann, S. M., Wise, M., Caylor, R., Imholt, F., Selimovic, V. and  
899 Shilling, J. E.: Optical properties and aging of light absorbing secondary organic aerosol, *Atmos. Chem. Phys.*, 16,  
900 12815–12827, doi:10.5194/acp-2016-482, 2016.
- 901 Liu, D. T., Whitehead, J., Alfarra, M. R., Reyes-Villegas, E., Spracklen, D. V., Reddington, C. L., Kong, S. F.,  
902 Williams, P. I., Ting, Y. C., Haslett, S., Taylor, J. W., Flynn, M. J., Morgan, W. T., McFiggans, G., Coe, H., and  
903 Allan, J. D.: Black-carbon absorption enhancement in the atmosphere determined by particle mixing state, *Nat.*  
904 *Geosci.*, 10, 184–188, doi:10.1038/ngeo2901, 2017.
- 905
- 906 Liu, P. F., Abdelmalki, N., Hung, H. M., Wang, Y., Brune, W. H. and Martin, S. T.: Ultraviolet and visible complex  
907 refractive indices of secondary organic material produced by photooxidation of the aromatic compounds toluene and  
908 m-xylene, *Atmos. Chem. Phys.*, 15(3), 1435–1446, doi:10.5194/acp-15-1435-2015, 2015a.
- 909 Liu, P. S. K., Deng, R., Smith, K. A., Jayne, J. T., Williams, L.R., Canagaratna, M. R., Moore, K., Onasch, T. B.,  
910 Worsnop, D.R., and Deshler, T.: Transmission efficiency of an aerodynamic focusing lens system: comparison of  
911 model calculations and laboratory measurements for the aerodyne aerosol mass spectrometer, *Aerosol Sci. Tech.*, 41,  
912 721–733, 2007.
- 913
- 914 Liu, S., Aiken, A. C., Gorkowski, K., Dubey, M. K., Cappa, C. D., Williams, L. R., Herndon, S. C., Massoli, P.,  
915 Fortner, E. C., Chhabra, P. S., Brooks, W. A., Onasch, T. B., Jayne, J. T., Worsnop, D. R., China, S., Sharma, N.,  
916 Mazzoleni, C., Xu, L., Ng, N. L., Liu, D., Allan, J. D., Lee, J. D., Fleming, Z. L., Mohr, C., Zotter, P., Szidat, S. and  
917 Prévôt, A. S. H.: Enhanced light absorption by mixed source black and brown carbon particles in UK winter, *Nat.*  
918 *Commun.*, 8435, doi:10.1038/ncomms9435, 2015b.
- 919 Lu, Z., Streets, D. G., Winijkul, E., Yan, F., Chen, Y., Bond, T. C., Feng, Y., Dubey, M. K., Liu, S., Pinto, J. P. and  
920 Carmichael, G. R.: Light absorption properties and radiative effects of primary organic aerosol emissions, *Environ.*  
921 *Sci. Technol.*, 49, 4868–4877, doi:10.1021/acs.est.5b00211, 2015.
- 922 Massabò, D., Caponi, L., Bernardoni, V., Bove, M. C., Brotto, P., Calzolari, G., Cassola, F., Chiari, M., Fedi, M. E.,  
923 Fermo, P., Giannoni, M., Lucarelli, F., Nava, S., Piazzalunga, A., Valli, G., Vecchi, R. and Prati, P.: Multi-  
924 wavelength optical determination of black and brown carbon in atmospheric aerosols, *Atmos. Environ.*, 108, 1–12,  
925 doi:10.1016/j.atmosenv.2015.02.058, 2015.
- 926 Massabò, D., Bernardoni, V., Bove, M. C., Brunengo, A., Cuccia, E., Piazzalunga, A., Prati, P., Valli, G. and  
927 Vecchi, R.: A multi-wavelength optical set-up for the characterization of carbonaceous particulate matter, *J. Aerosol*  
928 *Sci.*, 60, 34–46, doi:10.1016/j.jaerosci.2013.02.006, 2013.
- 929 Moosmüller, H., Chakrabarty, R. K. and Arnott, W. P.: Aerosol light absorption and its measurement: A review, *J.*  
930 *Quant. Spectrosc. Radiat. Transf.*, 110(11), 844–878, doi:10.1016/j.jqsrt.2009.02.035, 2009.
- 931 Moosmüller, H., Chakrabarty, R. K., Ehlers, K. M. and Arnott, W. P.: Absorption Ångström coefficient, brown  
932 carbon, and aerosols: Basic concepts, bulk matter, and spherical particles, *Atmos. Chem. Phys.*, 11(3), 1217–1225,  
933 doi:10.5194/acp-11-1217-2011, 2011.
- 934 Petzold, A. and Schönlinner, M.: Multi-angle absorption photometry - A new method for the measurement of  
935 aerosol light absorption and atmospheric black carbon, *J. Aerosol Sci.*, 35(4), 421–441,

- 936 doi:10.1016/j.jaerosci.2003.09.005, 2004.
- 937 Platt, S. M., El Haddad, I., Zardini, A. A., Clairotte, M., Astorga, C., Wolf, R., Slowik, J. G., Temime-Roussel, B.,  
938 Marchand, N., Ježek, I., Drinovec, L., Močnik, G., Möhler, O., Richter, R., Barmet, P., Bianchi, F., Baltensperger,  
939 U. and Prévôt, A. S. H.: Secondary organic aerosol formation from gasoline vehicle emissions in a new mobile  
940 environmental reaction chamber, *Atmos. Chem. Phys.*, 13(18), 9141–9158, doi:10.5194/acp-13-9141-2013, 2013.
- 941 Romonosky, D. E., Laskin, A., Laskin, J. and Nizkorodov, S. A.: High-resolution mass spectrometry and molecular  
942 characterization of aqueous photochemistry products of common types of secondary organic aerosols, *J. Phys.*  
943 *Chem. A*, 119(11), 2594–2606, doi:10.1021/jp509476r, 2015.
- 944 Saleh, R., Hennigan, C. J., McMeeking, G. R., Chuang, W. K., Robinson, E. S., Coe, H., Donahue, N. M. and  
945 Robinson, A. L.: Absorptivity of brown carbon in fresh and photo-chemically aged biomass-burning emissions,  
946 *Atmos. Chem. Phys.*, 13(15), 7683–7693, doi:10.5194/acp-13-7683-2013, 2013.
- 947 Saleh, R., Robinson, E. S., Tkacik, D. S., Ahern, A. T., Liu, S., Aiken, A. C., Sullivan, R. C., Presto, A. A., Dubey,  
948 M. K., Yokelson, R. J., Donahue, N. M., and Robinson, A. L.: Brownness of organics in aerosols from biomass  
949 burning linked to their black carbon content, *Nat. Geosci.*, 7, 2–5, doi:10.1038/ngeo2220, 2014.
- 950
- 951 Schauer, J. J., Mader, B. T., Deminter, J. T., Heidemann, G., Bae, M. S., Seinfeld, J. H., Flagan, R. C., Cary, R. A.,  
952 Smith, D., Huebert, B. J., Bertram, T., Howell, S., Kline, J. T., Quinn, P., Bates, T., Turpin, B., Lim, H. J., Yu, J. Z.,  
953 Yang, H., and Keywood, M. D.: ACE-Asia intercomparison of a thermal-optical method for the determination of  
954 particle-phase organic and elemental carbon, *Environ. Sci. Technol.*, 37, 993–1001, 2003.
- 955
- 956 Shafizadeh, F.: The chemistry of pyrolysis and combustion. The chemistry of solid Wood, *ACS Symp. Ser.*, 207,  
957 489–529, doi:10.1021/ba-1984-0207.ch013\|10.1021/ba-1984-0207.ch013, 1984.
- 958 Slowik, J. G., E. S. Cross, J.-H. Han, P. Davidovits, T. B. Onasch, J. T. Jayne, L. R. Williams, M. R. Canagaratna,  
959 D. R. Worsnop, R. K. Chakrabarty, H. Moosmüller, W. P. Arnott, J. P. Schwarz, R. S. Gao, D. W. Fahey, G. L. Kok  
960 and A. Petzold.: An inter-comparison of instruments measuring black carbon content of soot particles, *Aerosol Sci.*  
961 *Technol.* 41, 3, 295–314, 2007.
- 962
- 963 Sunlin, B. J., Pandey, A., Walker, M. J., Pattison, R. S., Williams, B. J., and Chakrabarty, R. K.: Atmospheric  
964 photooxidation diminishes light absorption by primary brown carbon aerosol from biomass burning, *Environ. Sci.*  
965 *Technol. Lett.*, 4(12), 540–545, doi:10.1021/acs.estlett.7b00393, 2017.
- 966
- 967 Sun, H., Biedermann, L. and Bond, T. C.: Color of brown carbon: A model for ultraviolet and visible light  
968 absorption by organic carbon aerosol, *Geophys. Res. Lett.*, 34(17), 1–5, doi:10.1029/2007GL029797, 2007.
- 969
- 970 Ulevicius, V., Bozzetti, C., Vlachou, A., Plauškaitė, K., Mordas, G., Dudoitis, V., Abbaszade, G., Remeikis, V.,  
971 Garbaras, A., Masalaite, A., Bles, J., Fröhlich, R., Dällenbach, K. R., Canonaco, F., Slowik, J. G., Dommen, J.,  
972 Zimmermann, R., Schnelle-kreis, J., Salazar, G. A. and Agrios, K., Szidat, S., Haddad, I. El., and Prévôt, A. S. H. :  
973 Fossil and non-fossil source contributions to atmospheric carbonaceous aerosols during extreme spring grassland  
974 fires in Eastern Europe, *Atmos. Chem. Phys.*, 16, 5513–5529, doi:10.5194/acp-16-5513-2016, 2016.
- 975
- 976 Wang, X., Heald, C. L., Ridley, D. A., Schwarz, J. P., Spackman, J. R., Perring, A. E., Coe, H., Liu, D. and Clarke,  
977 A. D.: Exploiting simultaneous observational constraints on mass and absorption to estimate the global direct  
978 radiative forcing of black carbon and brown carbon, *Atmos. Chem. Phys.*, 14(20), 10989–11010, doi:10.5194/acp-  
979 14-10989-2014, 2014.
- 980
- 981 Weingartner, E., Saathoff, H., Schnaiter, M., Streit, N., Bitnar, B. and Baltensperger, U.: Absorption of light by soot  
982 particles: determination of the absorption coefficient by means of aethalometers, *J. Aerosol Sci.*, 34(10), 1445–1463,  
doi:10.1016/S0021-8502(03)00359-8, 2003.
- 983
- 984 Yang, H. and Yu, J. Z.: Uncertainties in charring correction in the analysis of elemental and organic carbon in  
985 atmospheric particles by thermal/optical methods, *Environ. Sci. Technol.*, 36 (23), 5199–5204, 2002.

983 Zhang, X., Lin, Y. –H., Surratt, J. D., Zotter, P. and Prévôt, A. S. H.: Light-absorbing soluble organic aerosol in  
984 Los-Angeles and Atlanta: A contrast in secondary organic aerosol, *Geophys. Res. Lett.*, 38, 2011.  
985  
986 Zhao, R., Lee, A. K. Y., Huang, L., Li, X., Yang, F. and Abbatt, J. P. D.: Photochemical processing of aqueous  
987 atmospheric brown carbon, *Atmos. Chem. Phys.*, 15(11), 6087–6100, doi:10.5194/acp-15-6087-2015, 2015.

988 Zotter, P., Herich, H., Gysel, M., El-Haddad, I., Zhang, Y., Močnik, G., Hüglin, C., Baltensperger, U., Szidat, S. and  
989 Prévôt, A. S. H.: Evaluation of the absorption Ångström exponents for traffic and wood burning in the Aethalometer  
990 based source apportionment using radiocarbon measurements of ambient aerosol, *Atmos. Chem. Phys.*, 17, 4229-  
991 4249, doi:10.5194/acp-17-4229-2017, 2017.

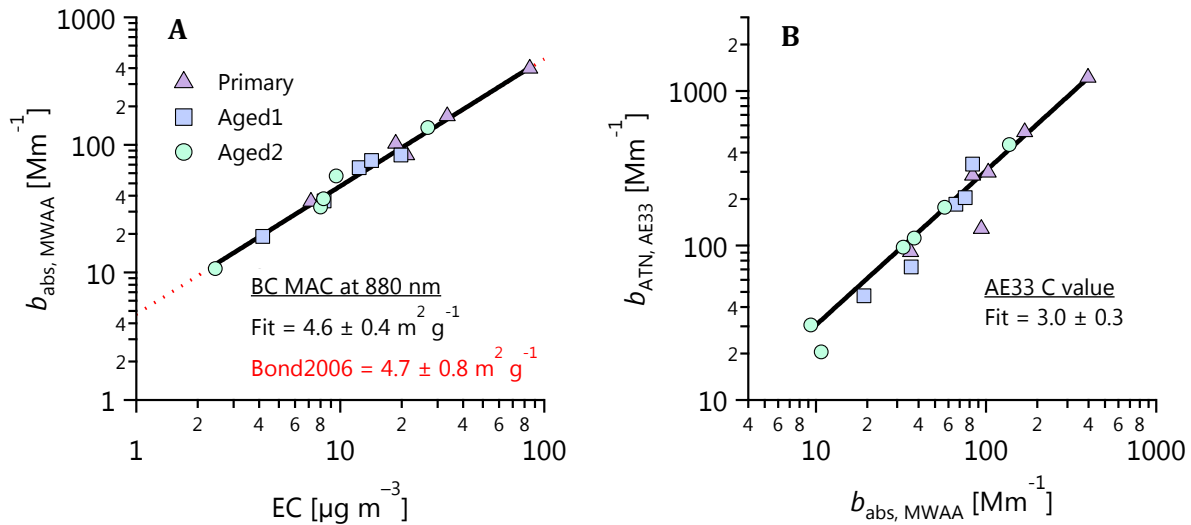
992 Zotter, P., Herich, H., Gysel, M., El-haddad, I., Zhang, Y. and Močnik, G.: Evaluation of the absorption Ångström  
993 exponents for traffic and wood burning in the Aethalometer-based source apportionment using radiocarbon  
994 measurements of ambient aerosol, , 4229–4249, doi:10.5194/acp-17-4229-2017, 2017.

995  
996  
997  
998  
999  
1000  
1001  
1002  
1003  
1004  
1005  
1006  
1007  
1008  
1009  
1010  
1011  
1012  
1013  
1014  
1015  
1016  
1017



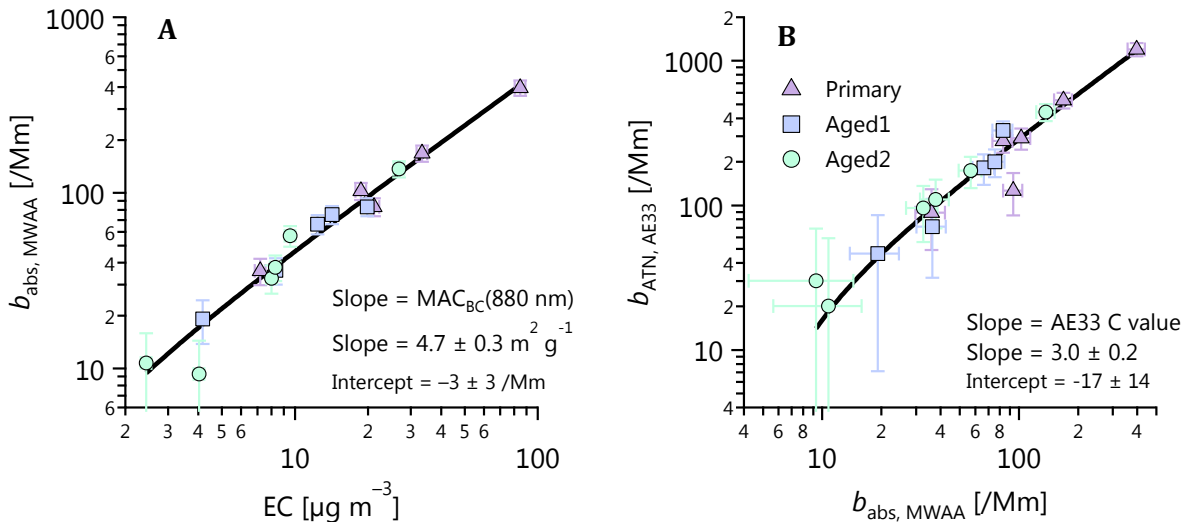
1018

1019 **Old Figure**



1020

1021 **New Figure**



1022

1023 **Figure 1: Determination of (A)  $MAC_{BC}(880nm)$  and (B) aethalometer C value using MWAAs absorption measurements,**  
 1024 **thermal/optical EC (EUSAAR2 protocol) and aethalometer attenuation measurements. MWAAs absorption measurements**  
 1025 **at 880 nm is determined by extrapolating the absorption coefficients at 850 nm using an  $\alpha$  determined from the ratio**  
 1026 **between the absorption coefficients at 850 nm and 635nm. The aerosols were either primary (no OH exposure), Aged 1**  
 1027 **( $\sim 1 \times 10^7$  molec OH  $cm^{-3}$  h), or Aged 2 ( $\sim 4 \times 10^7$  molec OH  $cm^{-3}$  h). No difference in MAC or C value was discernable with**  
 1028 **aging (see also Fig. S2). Also shown is the MAC of pure BC recommended by Bond et al. (2006) (dotted line in A). The C**  
 1029 **value derived from  $\sigma_{ATN}$  recommended by Drinovec et al. (2015) = 2.6 compares well with the value derived in Fig. 1B.**

1030

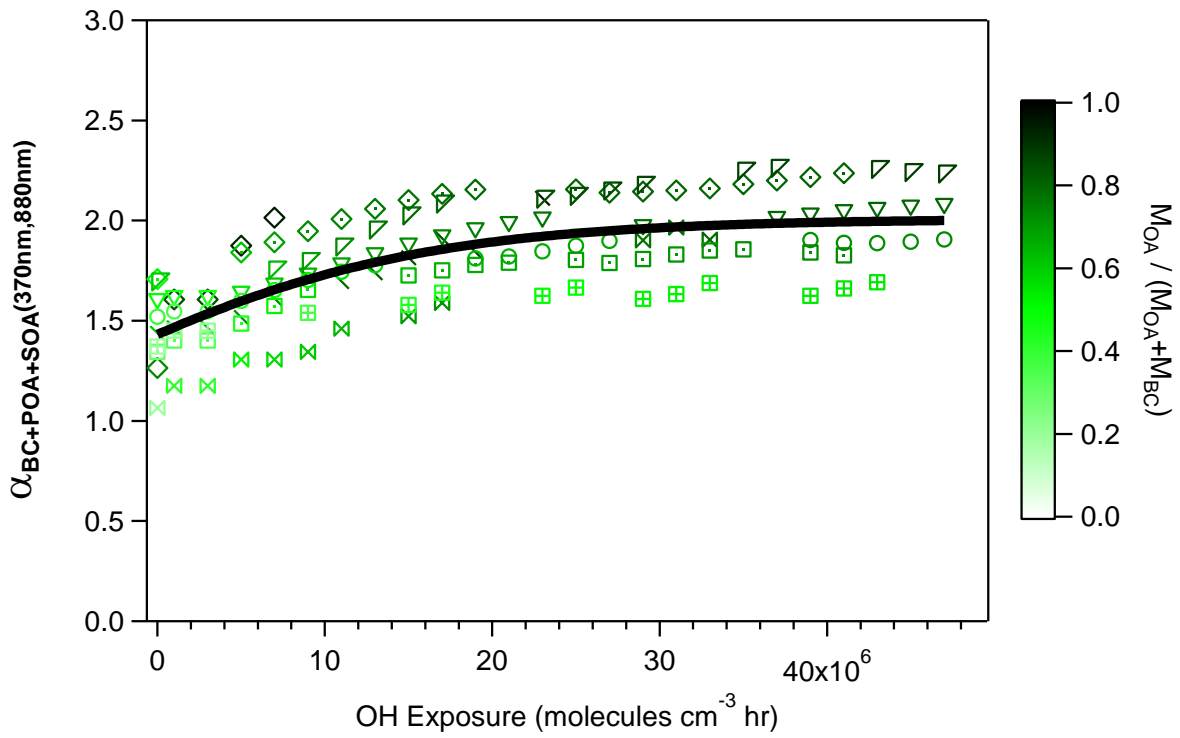
1031

1032

1033

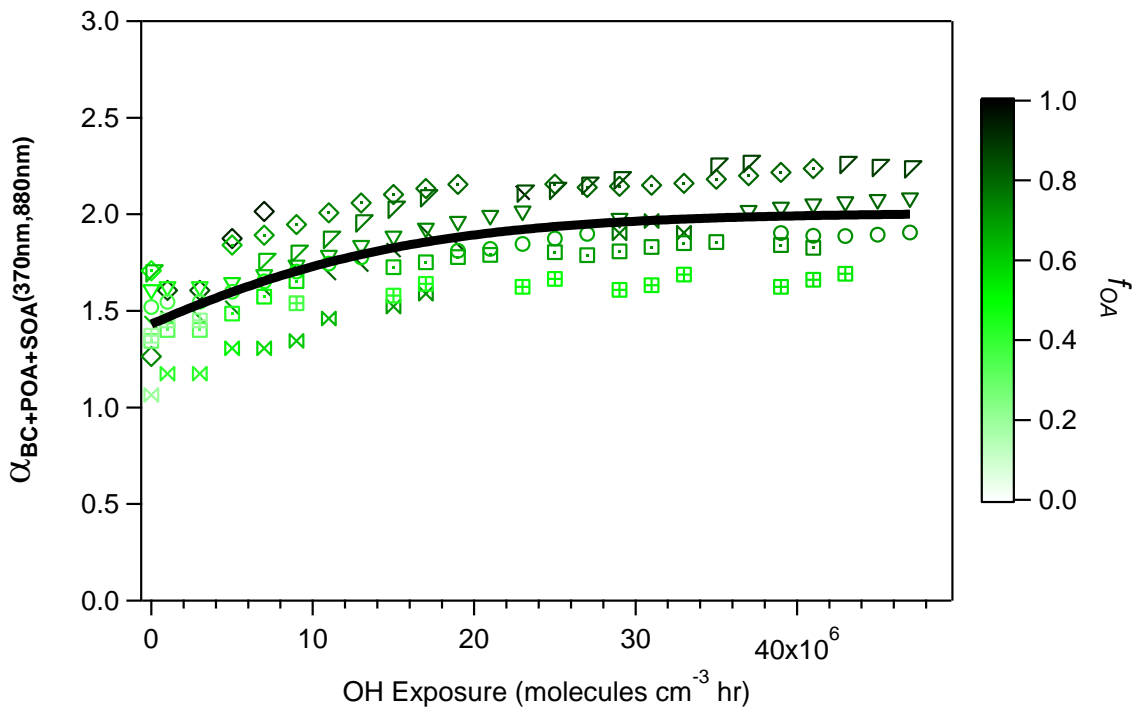
1034

1035 **Old Figure**



1036

1037 **New Figure:**



1038

1039 Figure 2: Evolution during photochemical aging of  $\alpha_{BC+POA+SOA}(370nm, 880nm)$  (two-wavelength Ångström exponent  
 1040 calculated using total absorption data at 370 nm and 880 nm), where the different symbols denote individual experiments.  
 1041 Data are colored by the OA mass fraction  $f_{OA} = M_{OA}/(M_{OA} + M_{BC})$ . The black line is a fit to guide the eye.

1042

1043 Old Figure:

1044

1045

1046

1047

1048

1049

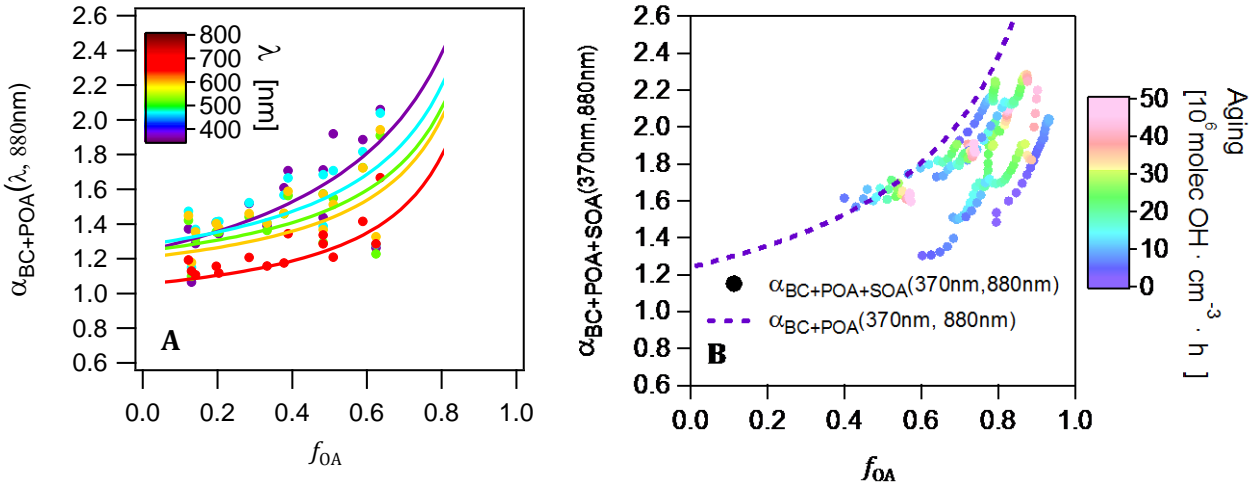
1050

1051

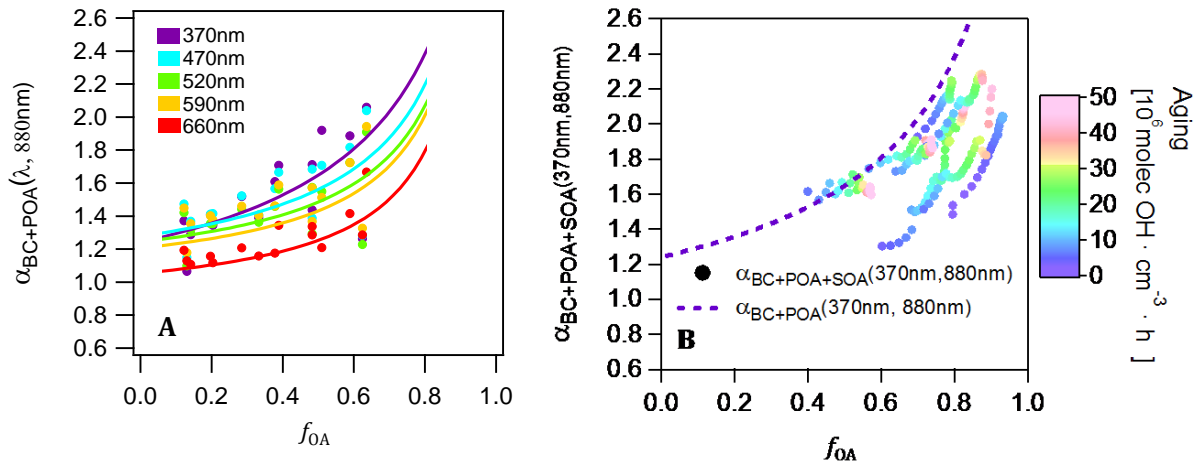
1052

1053

1054



1054 New Figure:



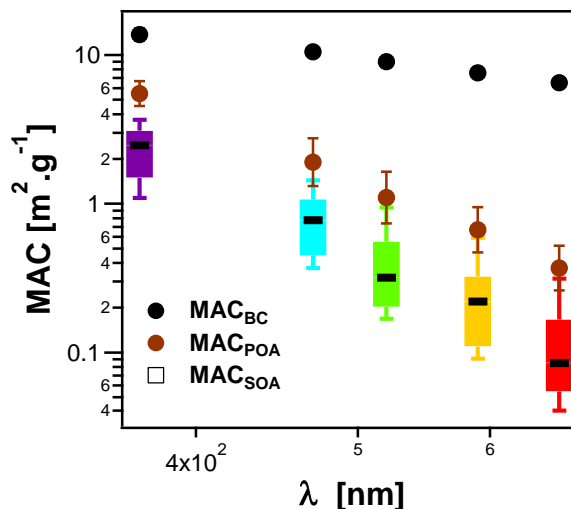
1055

1056 **Figure 3: (A) Relationship of  $\alpha_{BC+POA}(\lambda, 880nm)$  to  $f_{OA}$  for seven wavelengths. Lines are fits of Eq. (13) to the data. (B)**  
 1057 **Relationship of  $\alpha_{BC+POA+SOA}(370nm, 880nm)$  to  $f_{OA}$  for several experiments. Data in (A) and (B) are colored by the**  
 1058 **wavelength and OH exposure, respectively.**

1059 **Figure 3: (A) Relationship of  $\alpha_{BC+POA}(\lambda, 880nm)$  to  $f_{OA}$  for seven wavelengths for primary emissions. Data are colored**  
 1060 **by the wavelength. Curves are fits of Eq. (13) to the data. Each point represents the average of one experiment and**  
 1061 **therefore the variability in  $f_{OA}$  is related to the variability in the emission composition between experiments. (B)**  
 1062 **Relationship of  $\alpha_{BC+POA+SOA}(370nm, 880nm)$  to  $f_{OA}$  for several experiments for aged aerosols. Data are color coded by**  
 1063 **the OH exposure. The variability in  $f_{OA}$  is due to SOA formation with aging; data from several experiments are shown**  
 1064 **which explains the wide range of  $f_{OA}$  at low OH exposures. Note that more data are included in A than B, as primary**  
 1065 **emissions for some experiments were not aged.**

1066

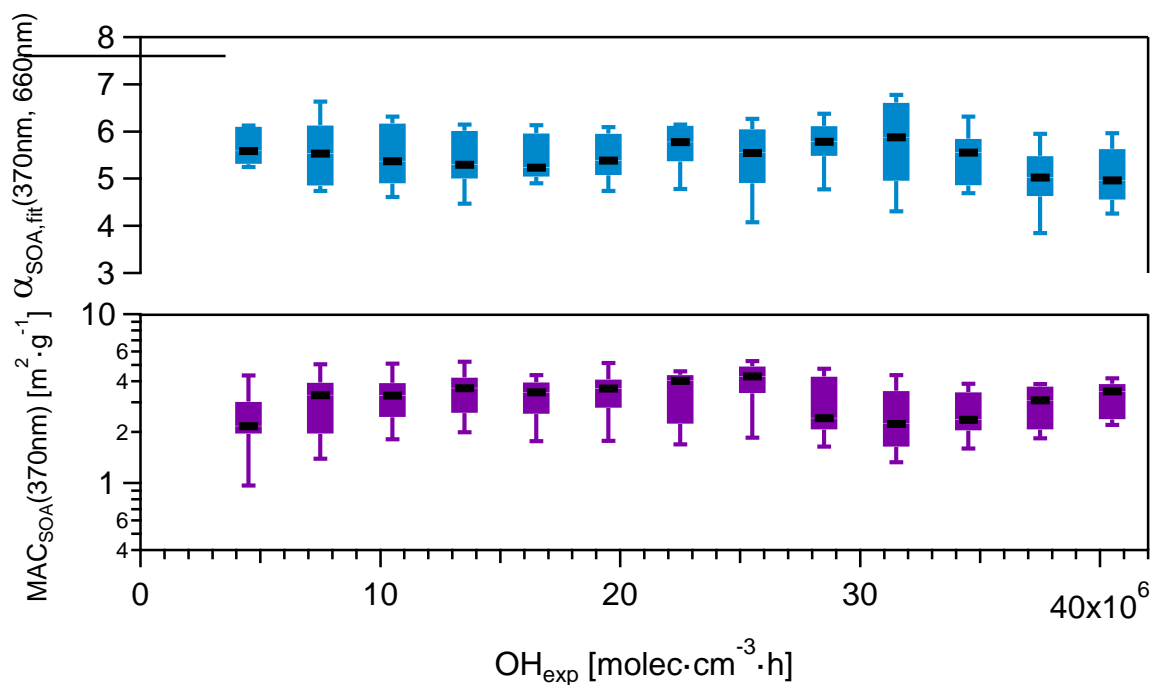
1067



1068

1069 Figure 4:  $MAC_{SOA}(\lambda)$  calculated from several smog chamber experiments plotted as box-whiskers as a function of  
 1070 wavelength (also shown by the color of the bars). The thick black lines, the boxes and the whiskers mark the medians, the  
 1071 quartiles and the 10<sup>th</sup> and the 90<sup>th</sup> percentiles, respectively. Also shown are the  $MAC_{BC}(\lambda)$  and  $MAC_{POA}(\lambda)$  reported in  
 1072 Table 1. Note that  $MAC_{SOA}(880nm)$  and  $MAC_{POA}(880nm)$  are zero by definition.

1073



1074

1075 Figure 5:  $MAC_{SOA}(370nm)$  and  $\alpha_{SOA,fit}(370nm, 660nm)$  calculated from several smog chamber experiments plotted as a  
 1076 function of OH exposure. The box marks the 25<sup>th</sup> and 75<sup>th</sup> percentile, while the whiskers mark the 10<sup>th</sup> and the 90<sup>th</sup>  
 1077 percentile.  $MAC_{SOA}(370nm)$  was obtained using Eq. (19).  $\alpha_{SOA,fit}(370nm, 660nm)$  was obtained from fitting the  $MAC_{SOA}$   
 1078 values in the range 370-660 nm for the different experiments against the wavelength.  $\alpha_{SOA,fit}(370nm, 660nm)$  is the slope  
 1079 of the linear fit applied after log transforming the data.  $MAC_{SOA}(\lambda)$  for higher wavelengths are shown in Fig. S10.

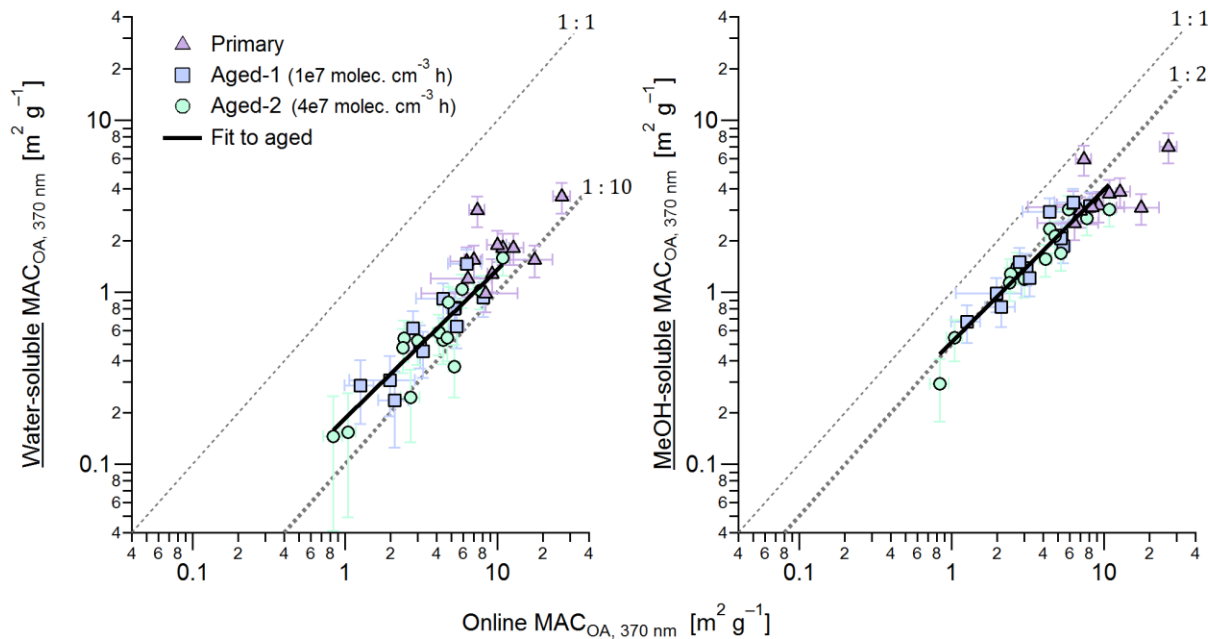
1080

1081

1082

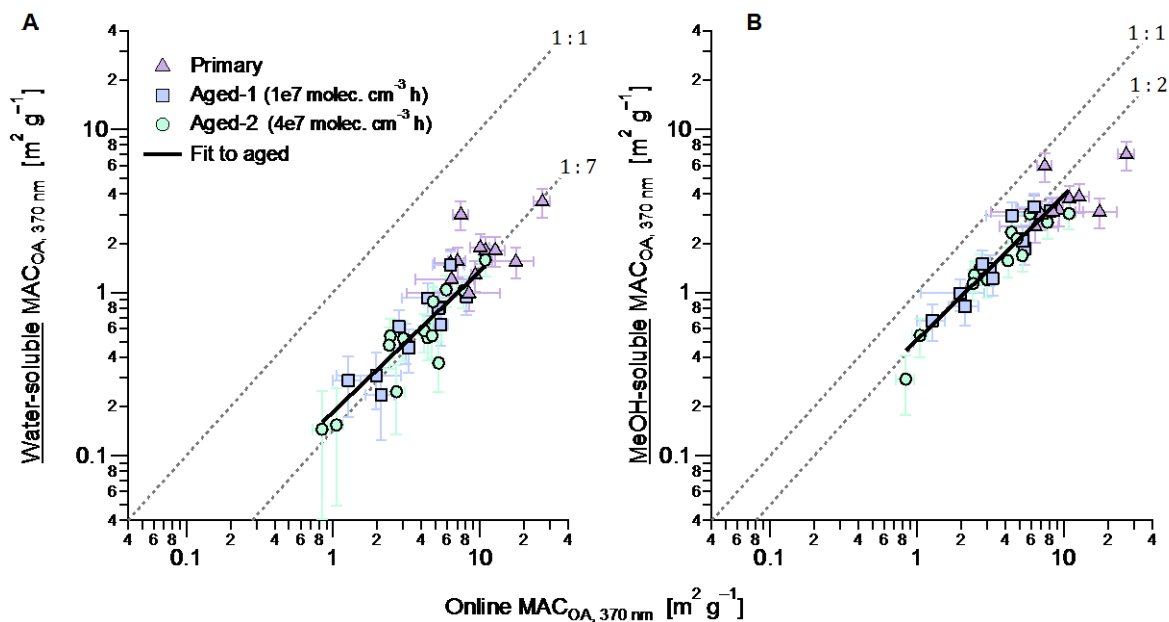
1083

1084 **Old Figure:**



1085

1086 **New Figure:**



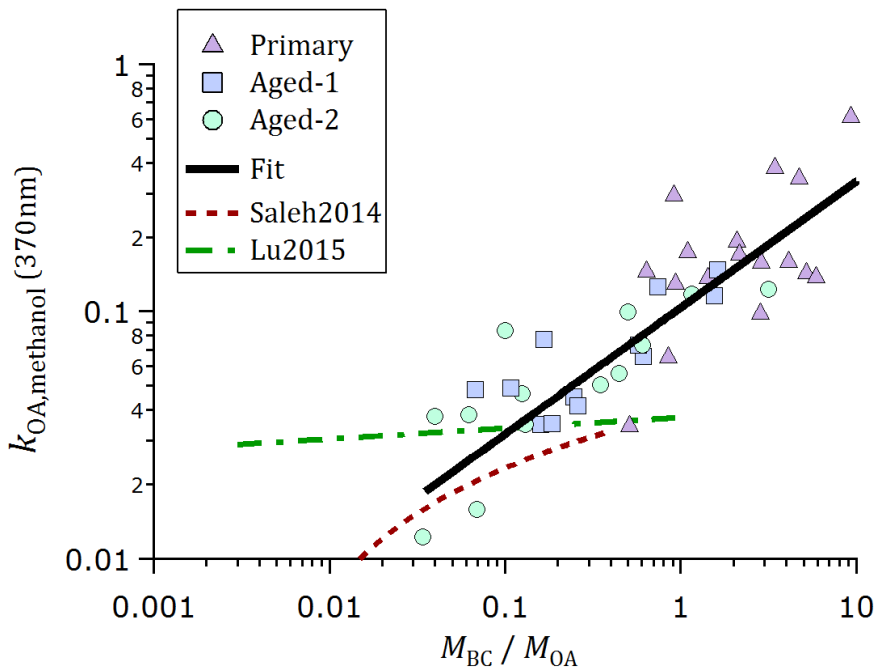
1087

1088 **Figure 6: Comparison of the MAC<sub>OA</sub>(370nm) of aged aerosols determined from online and offline measurements of**  
1089 **absorption. The offline filter extraction method directly quantified properties of total OA (ordinate), while the average of**  
1090 **MAC<sub>SOA</sub> and MAC<sub>POA</sub> from the online measurements weighted with respective mass concentrations is shown on the**  
1091 **abscissa. (A) offline measurements of water-soluble OA, (B) methanol-soluble OA.**

1092 **Figure 6: Comparison of the MAC<sub>OA</sub>(370nm) of aged aerosols determined from online and offline absorption**  
1093 **measurements. The offline filter-extraction method directly quantified properties of total OA (ordinate), while the**  
1094 **average of MAC<sub>SOA</sub> and MAC<sub>POA</sub> weighted with respective mass concentrations is shown on the abscissa. The panels show**  
1095 **offline measurements of (A) water-soluble OA, (B) methanol-soluble OA. Fits are to aged data only due to the**

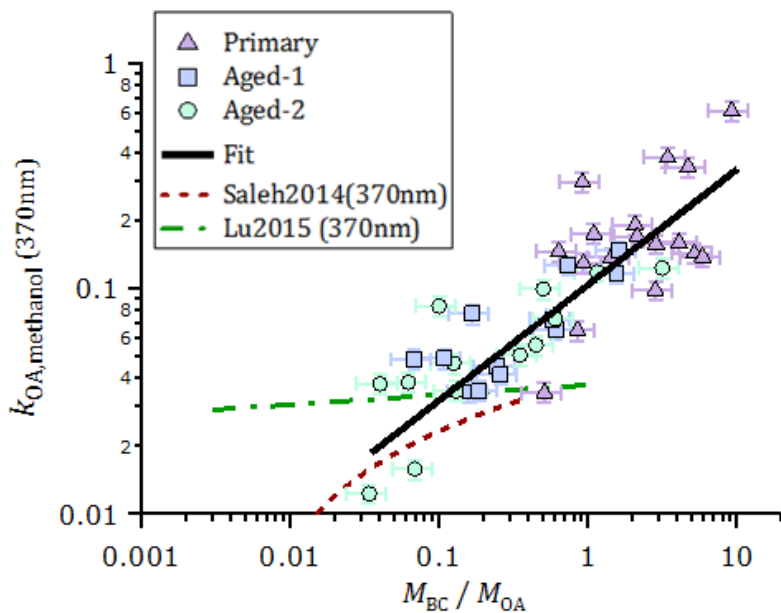
1096 significantly smaller scatter of those data, although primary data on average follow similar trends. The fitted slopes and  
1097 intercepts are, respectively, (A)  $0.13 \pm 0.02$  and  $0.05 \pm 0.06 \text{ m}^2 \text{ g}^{-1}$  and (B)  $0.12 \pm 0.1$  and  $0.38 \pm 0.03 \text{ m}^2 \text{ g}^{-1}$ .  
1098

1099 Old Figure



1100

1101 New Figure



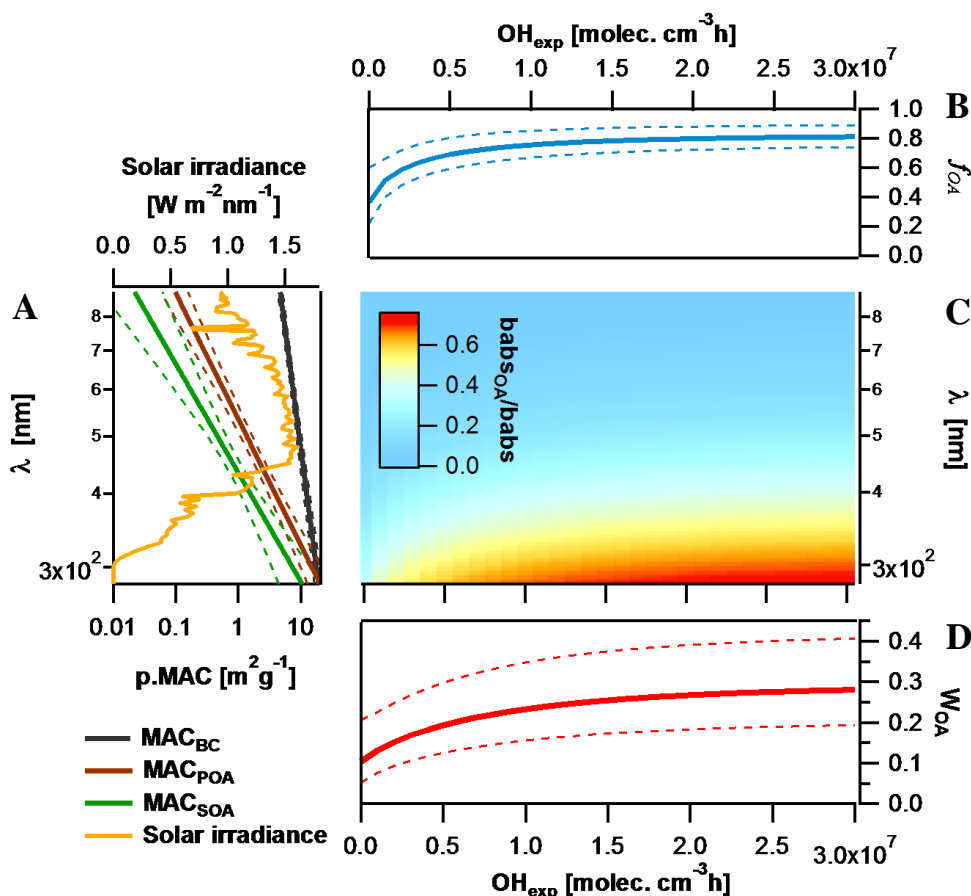
1102

1103 **Figure 7: Imaginary part of the OA refractive index at 370 nm, obtained from offline UV/vis spectroscopy of methanol**  
 1104 **OA extracts, plotted as a function of  $f_{OA}$ . The ordinary least-squares fit is  $\log(k_{OA,370}) = \log(M_{BC}/M_{OA})(0.51 \pm$   
 1105  **$0.07) + (-0.98 \pm 0.05)$ .****

1106 **Figure 7: Imaginary part of the OA refractive index at 370 nm, obtained from offline UV/vis spectroscopy of methanol**  
 1107 **OA extracts, plotted as a function of  $f_{OA}$ . The data could be empirically represented by a linear function in the log-log**  
 1108 **space, in the measurement range. The ordinary least-squares fit is  $(k_{OA, nm}) = \log(M_{BC}/M_{OA})(0.51 \pm 0.07) + (-0.98 \pm 0.05)$ .**  
 1109 **Also shown are parameterizations of  $k_{OA}(370 \text{ nm})$  for open burning against  $M_{BC}/M_{OA}$  estimated based on the online  $k_{OA}$**   
 1110 **(550 nm) measurements in Saleh et al. (2014) and Lu et al. (2015), using the  $k_{OA}$  wavelength dependence reported by the**  
 1111 **respective authors.**  
 1112

1113

1114



1115

1116 **Figure 8: Impact of BrC absorption on total primary and secondary wood-burning-aerosol absorption. (A) MACs of**  
 1117 **different particle components (BC, POA and SOA) along with their corresponding standard deviations plotted as a**  
 1118 **function of wavelength based on smog chamber data and compared to the solar irradiance spectrum. (B) Species average**  
 1119 **relative abundance in the smog chamber ( $f_{OA}$ ) plotted as a function of the OH exposure. (C) Image plot showing the OA**  
 1120 **absorption coefficient relative to the total aerosol absorption as a function of wavelength and OH exposure. (D) Rate of**  
 1121 **energy transfer due to BrC light absorption relative to the total carbonaceous aerosol absorption ( $W_{OA}$ ) estimated as a**  
 1122 **function of aging using the solar flux, the fractions of the different components and their MACs.**

1123

1124

1125

1126 **Table 1: Geometric mean and standard deviations of the determined MACs of BC, POA and SOA at different**  
1127 **wavelengths. Uncertainties were obtained from fits of Eq. (13) for  $MAC_{BC}$ ,  $MAC_{POA}$ , while for  $MAC_{SOA}$  uncertainties**  
1128 **GSD values are geometric standard deviation values on the  $MAC_{SOA}$  average values from all experiments. These**  
1129 **uncertainties do not include uncertainties related to the determination of  $MAC_{BC}(880nm)$ . By definition, BrC absorbance**  
1130 **at 880 nm is zero.**

$\lambda$ (nm)	BC		POA		SOA	
	GM ( $m^2 g^{-1}$ )	GSD	GM ( $m^2 g^{-1}$ )	GSD	GM ( $m^2 g^{-1}$ )	GSD
370	13.7	1.1	5.5	1.21	2.2	1.39
470	10.5	1.06	1.9	1.45	0.72	1.61
520	9	1.04	1.1	1.49	0.34	1.75
590	7.6	1.03	0.67	1.42	0.2	1.97
660	6.5	1.01	0.37	1.41	0.09	2.42
880	4.6	0.7	0*		0*	

1131

\*By definition

1132



# Supporting information for:

## Production of particulate brown carbon during atmospheric aging of **residential** wood-burning emissions

Nivedita K. Kumar<sup>1</sup>, Joel C. Corbin<sup>1\*</sup>, Emily A. Bruns<sup>1</sup>, Dario Massabó<sup>2</sup>, Jay G. Slowik<sup>1</sup>, Luka Drinovec<sup>3,4</sup>, Griša Močnik<sup>3,4</sup>, Paolo Prati<sup>2</sup>, Athanasia Vlachou<sup>1</sup>, Urs Baltensperger<sup>1</sup>, Martin Gysel<sup>1</sup>, Imad El-Haddad<sup>1</sup> and André S. H. Prévôt<sup>1</sup>

<sup>1</sup>Laboratory of Atmospheric Chemistry, Paul Scherrer Institute, 5232 Villigen, Switzerland

<sup>2</sup>Department of Physics & INFN, University of Genoa, via Dodecaneso 33, 16146, Genova, Italy

<sup>3</sup>Aerosol d.o.o, Kamniška 41, 1000 Ljubljana, Slovenia

<sup>4</sup>Condensed Matter Physics, Jožef Stefan Institute, 1000 Ljubljana, Slovenia

\*Now at National Research Council Canada, Ottawa, Canada

Correspondence to: I. El-Haddad ([imad.el-haddad@psi.ch](mailto:imad.el-haddad@psi.ch)), A. S. H. Prévôt ([andre.prevot@psi.ch](mailto:andre.prevot@psi.ch))

### Wall loss correction

Solving the equations in Section 3.1 requires the determination of the time-dependent concentrations of the different absorbing species, which may be governed by their photochemical production or decay as well as by diffusion, electrostatic and gravitational losses to the walls. Assuming all particles are equally lost to the walls, an inert, non-volatile species,  $X$ , follows a first order decay:

$$X(t) = X(t_0) \cdot \exp(\tau^{-1}(t - t_0)) \quad (\text{S1})$$

Here,  $t$  and  $t_0$  denote the time of interest and reference time, respectively. The time constant  $\tau$  is the lifetime of  $X$  with respect to particle wall losses. We determined  $\tau$  by fitting  $b_{\text{abs}}(t, 880\text{nm})$  to Equation S1. Only the last period of each experiment was chosen for fitting, when secondary organic aerosol production rates are smaller. On average,  $\tau$  equals  $3.9 \pm 0.8$  hours for our chamber. The wall loss corrected absorption coefficient,  $b_{\text{abs}}^{\text{WLC}}(t, 880\text{nm})$ , varied less than 8% throughout the experiment, with higher values in the first period of

measurements. Therefore we conclude that the first order decay is an appropriate approach for the wall loss correction of inert particulate properties. We ascribe the residual variations of  $b_{\text{abs}}^{\text{WLC}}(t, 880 \text{ nm})$  to a combination of uncertainties, including the aethalometer compensation parameter and possible small changes of  $\text{MAC}_{\text{BC}}(880\text{nm})$  with aging.

For the extrapolation of our data to ambient environments we computed the average SOA mass formed as a function of OH exposure during the different experiments. This step requires the correction of OA mass for particle wall losses, which has been achieved by assuming two cases: (1) condensable oxidized gases do not interact with wall-deposited particles and (2) condensable oxidized gases condense at similar rates onto the suspended and wall-deposited particles (Pierce et al., 2008). We did not consider the deposition of oxidized vapors onto the clean Teflon walls, which would require knowledge of the saturation vapor pressures of the compounds, the condensed phase bulk properties and the vapor-wall equilibration rates. It is likely that the large particle condensational sinks utilized here (with a particle surface area concentration of several hundreds of  $\mu\text{m}^2 \text{ cm}^{-3}$ ), outcompeted vapor deposition onto the walls. Therefore, we consider the vapor deposition to the clean Teflon wall to be of a minor importance compared to burn-to-burn variability and other experimental uncertainties.

Solving the mass balance equations of the suspended organic aerosol,  $[OA_{\text{sus}}(t)]$ , and the organic aerosol on the walls yields the expressions in Equations (S2) and (S3), when considering scenario (1) and (2), respectively:

$$[M_{OA, \text{wlc}, 1}(t)] = [M_{OA, \text{sus}}(t)] + \int_0^t \tau^{-1} [M_{OA, \text{sus}}(t)] dt \quad (\text{S2})$$

$$[M_{OA, \text{wlc}, 2}(t)] = [M_{OA, \text{sus}}(t)] * \exp(\tau^{-1} t) \quad (\text{S3})$$

Here,  $[M_{OA, \text{wlc}}(t)]$  is the wall-loss-corrected OA concentration. The results presented in Fig. 8 in the manuscript are the average time-series of all experiments considering both scenarios, and associated ranges entail both the experiment-to-experiment variability and the uncertainties related to wall loss corrections.

# Supplementary figures

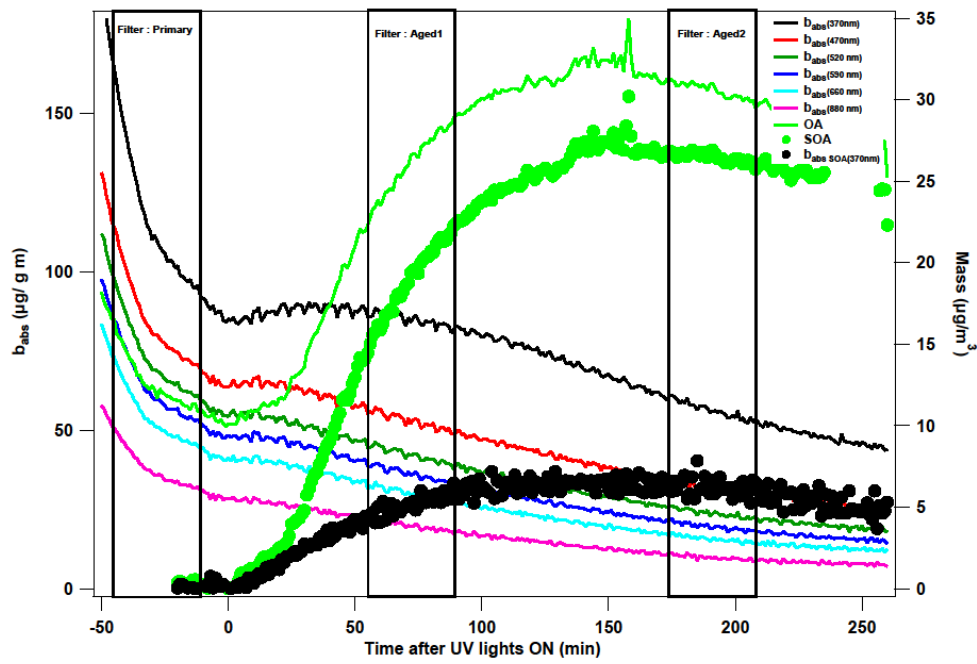
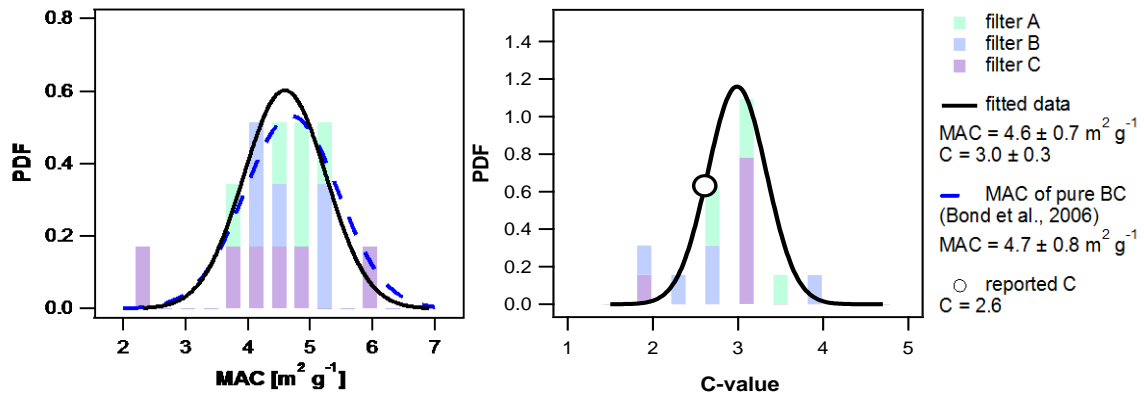


Figure S1: Absorption coefficients of fresh and aged emissions measured at 6 different wavelengths (i.e. 370 – 880 nm) using the aethalometer. The OA is measured using an AMS. Dotted lines are primary-subtracted OA (SOA) and absorption coefficient ( $b_{\text{absSOA}}(370\text{nm})$ ). The black boxes mark the times where the primary, slightly aged (Aged1, OH exposure  $\sim 1 \times 10^7$  molecules  $\text{cm}^{-3}$  h) and heavily aged filters (Aged2, OH exposure  $\sim 4 \times 10^7$  molecules  $\text{cm}^{-3}$  h) were collected.

## Old Figure:



New Figure:

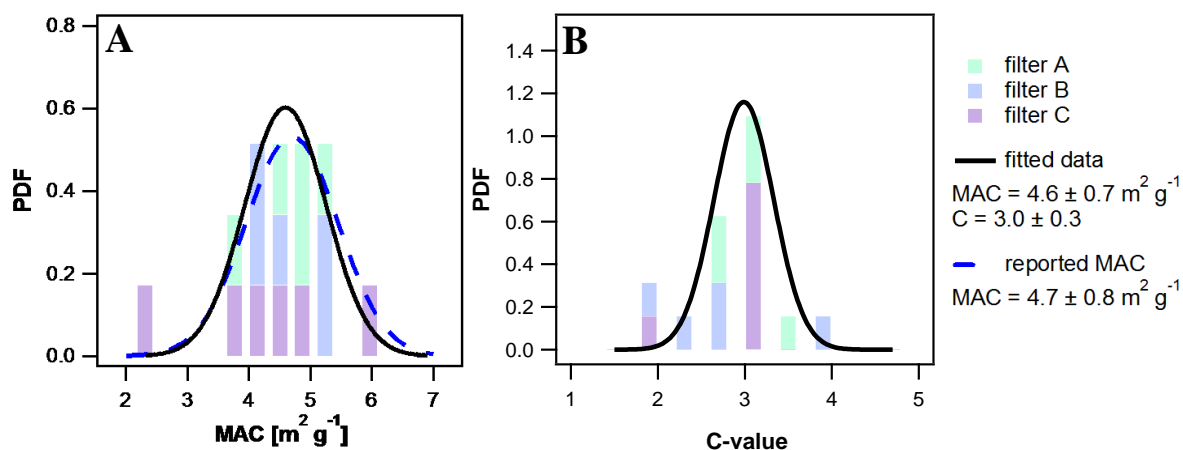


Figure S2: (a) Probability density function (PDF) comparing the MAC values determined by normalizing MWAA absorption measurements of offline primary (filter A), slightly aged (filter B: Aged1) and aged filter (filter C: Aged2) samples to EC (EUSAAR2) measurements of the same samples (bold line). A literature value for pure BC is also shown (Bond et al., 2006) (dashed blue line). (b) PDF comparing aethalometer attenuation measurements at 880 nm and MWAA absorption measurements at 850 nm to retrieve the aethalometer C value.

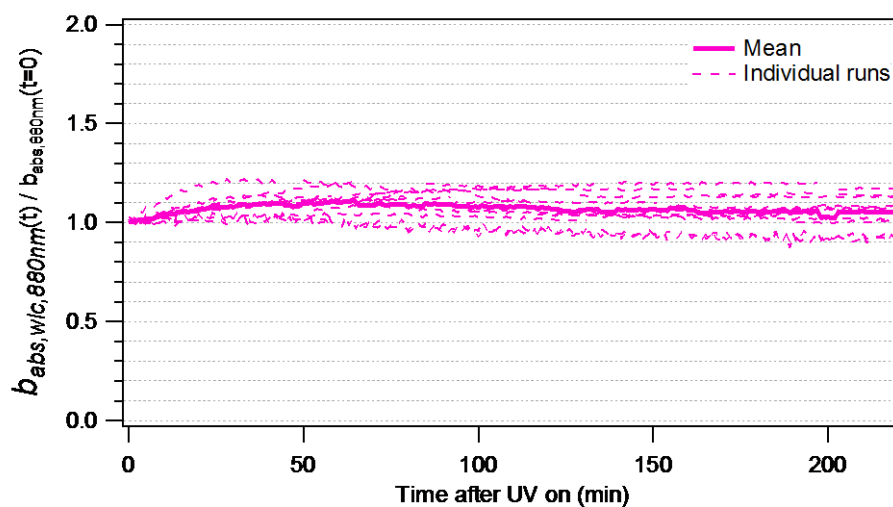
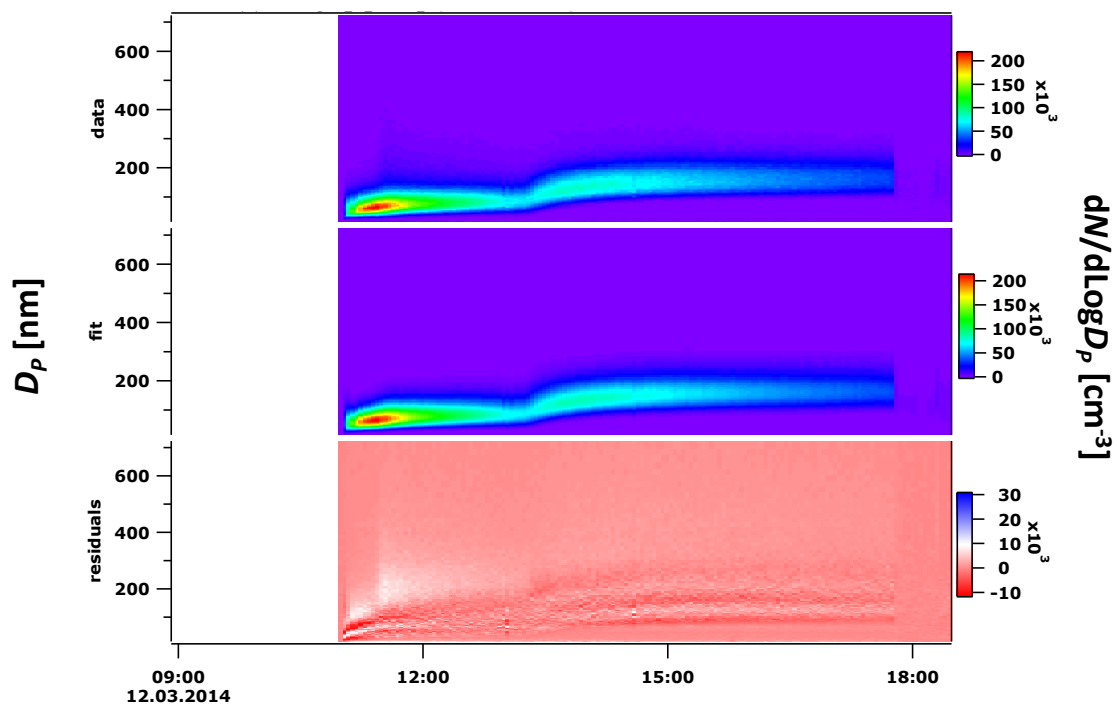


Figure S3: Wall-loss-corrected aethalometer absorption coefficient at 880 nm normalized to start-of-experiment absorption. The lack of any trend in this plot illustrates that the wall loss correction is appropriate and that only a negligible absorption increase occurs due to additional lensing by SOA.

Old Figure:



New Figure:

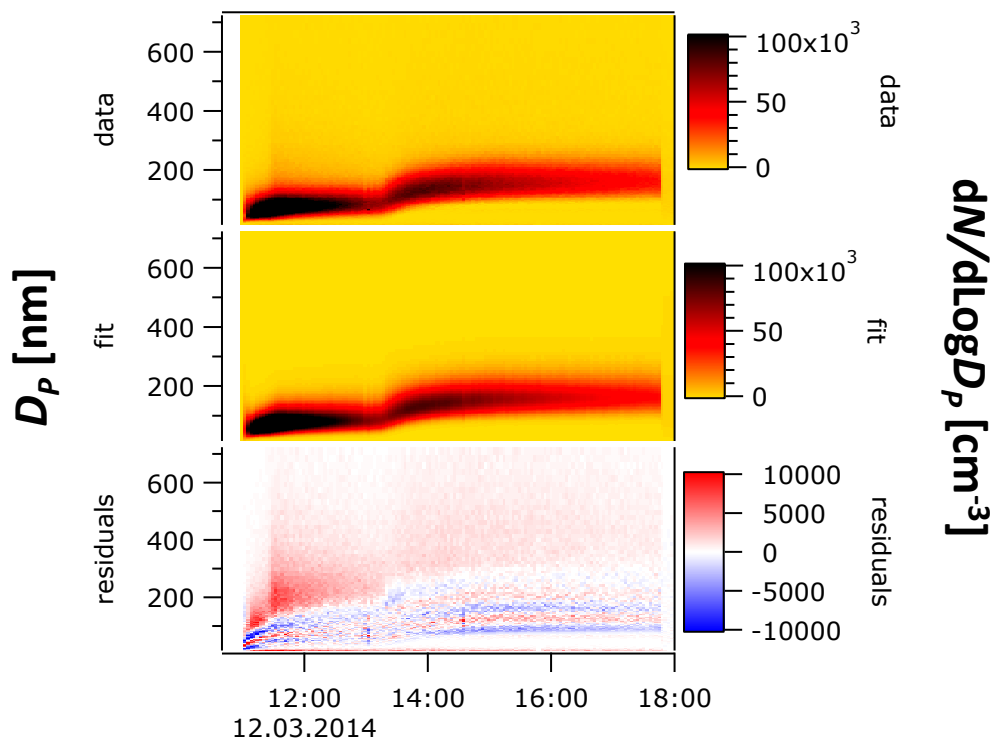
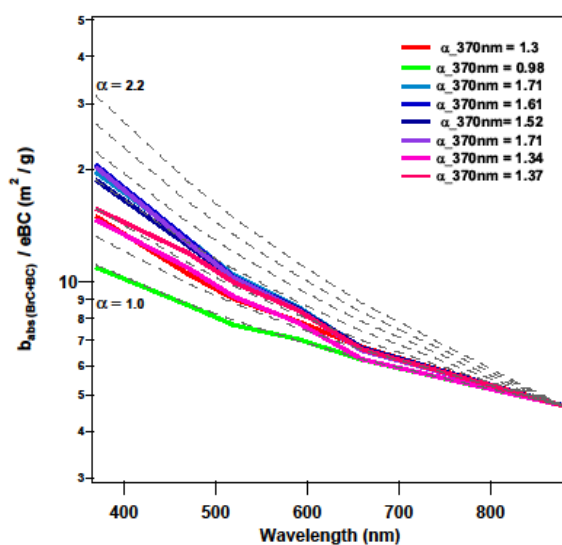


Figure S4: SMPS measurements (top), lognormal fits (middle;  $f_{\text{POA}}=0.51$ ), and fit residuals (bottom) of the size distribution of biomass burning organic aerosol during a typical aging experiment.

Old Figure:



New Figure:

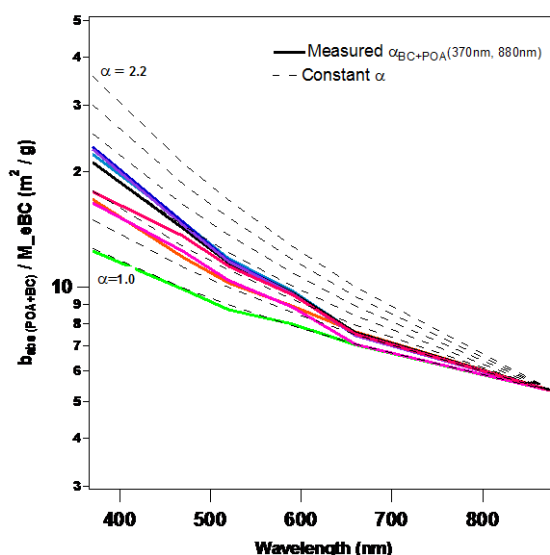
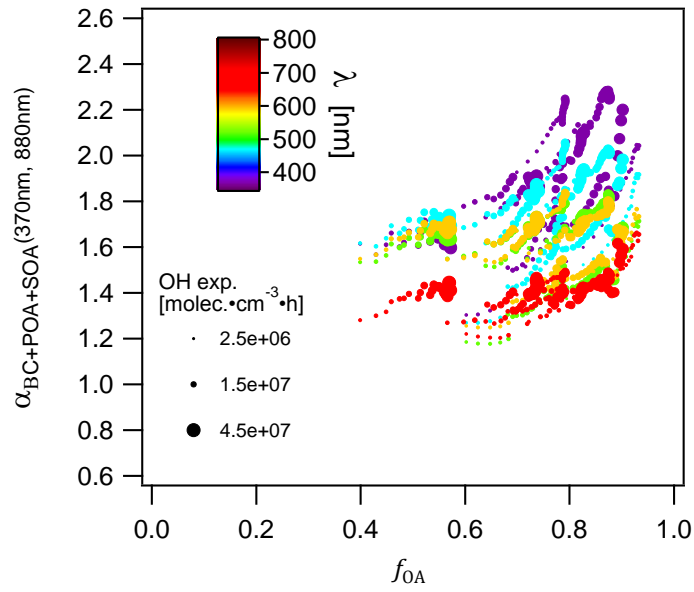


Figure S5: Absorption coefficients of fresh wood burning emissions measured using an aethalometer normalized to the eBC mass as a function of wavelength. In the legend each color denotes the  $\alpha_{\text{BC+POA}}(370\text{nm}, 880\text{nm})$  for an individual experiment. The dashed lines mark the literature range of  $\alpha$  values obtained for primary biomass burning emissions.

Figure S5: Absorption coefficients of fresh wood burning emissions measured using an aethalometer normalized to the eBC mass as a function of wavelength. In the legend each color denotes the  $\alpha_{\text{BC+POA}}(370\text{nm}, 880\text{nm})$  for an individual experiment. The dashed lines mark the absorption profiles calculated assuming a constant  $\alpha_{\text{BC+POA}}$  in the range 370-880nm. The range of  $\alpha$  values, is set between 1-2.2 (with an increment of 0.2), based on literature reports for primary biomass burning emissions from residential heating. The observed absorption spectra have steeper gradients with decreasing wavelength compared to the lines of constant alpha. This systematic decrease in  $\alpha(\lambda, 880\text{nm})$  with increasing  $\lambda$  reflects the more-efficient light absorption by BrC at shorter wavelengths (Moosmüller et al., 2011), and shows that the power law wavelength dependence is an inaccurate oversimplification for this mixed aerosol.

Old Figure:



New Figure:

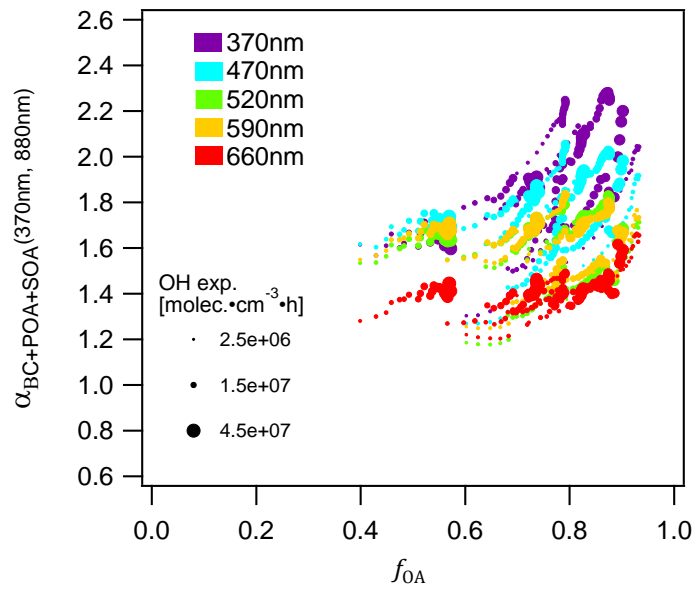
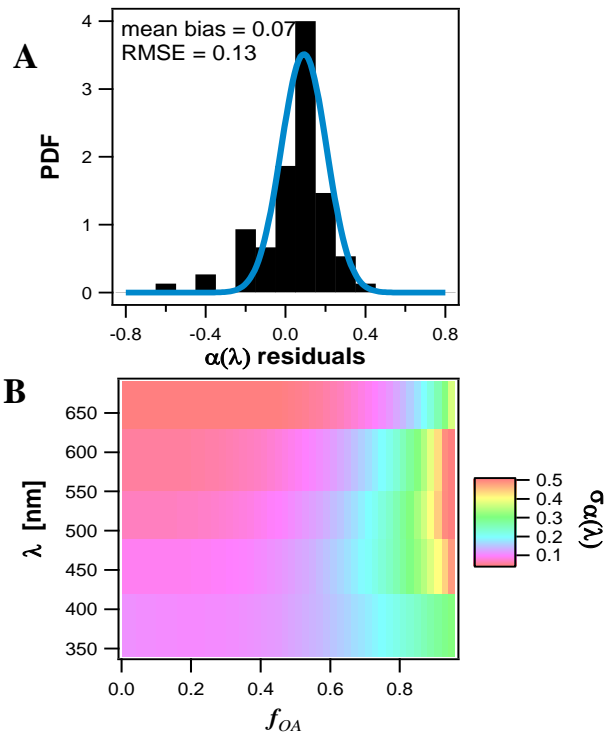


Figure S6: Relationship of  $\alpha_{BC+POA+SOA}(\lambda, 880nm)$  to  $f_{OA}$  for seven wavelengths, with symbol sizes indicating OH exposure.

Old Figure:



New Figure:

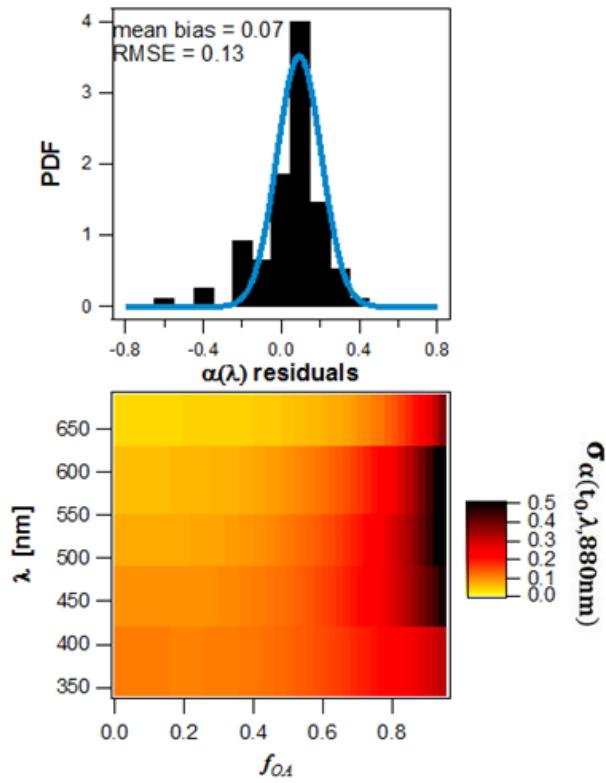




Figure S7: Analysis of the fitting errors of  $\alpha(\lambda, 880\text{nm})$  of primary emissions as a function of  $f_{OA}$ . Panel A shows the  $\alpha$  residual as a probability density function. Panel B is an image plot of the  $\alpha(\lambda, 880\text{nm})$  uncertainty as a function of  $f_{OA}$  estimated based on error propagation, using Equation 13 and assuming  $MAC_{POA}(\lambda)$ ,  $MAC_{BC}(\lambda)$  and  $f_{OA}$  to be independent variables (assuming negligible covariance between these quantities). The image plot in panel B shows that at short wavelengths and low fractions of OA, the estimated error on  $\alpha$  is less than 0.1. However, with increasing OA fraction and at longer wavelength the uncertainty increases. It should be noted that this increasing uncertainty is not greater than 0.5 even at OA fractions of 0.99, corresponding to an  $\alpha$  of  $\sim 2.2$ .

Figure S7: Analysis of the fitting errors of  $\alpha(\lambda, 880\text{nm})$  of primary emissions as a function of  $f_{OA}$ . Panel A shows the  $\alpha$  residual as a probability density function. Panel B is an image plot of the  $\alpha(\lambda, 880\text{nm})$  error,  $\sigma_{\alpha(t_0, \lambda, 880\text{nm})}$ , as a function of  $f_{OA}$  at different wavelengths.  $\sigma_{\alpha(t_0, \lambda, 880\text{nm})}$  is obtained from the error propagation of Eq. (13) solved for different wavelengths, using the geometric mean and standard deviation of  $MAC_{POA}(\lambda)$  and  $MAC_{BC}(\lambda)$ . This error term represents the variability in or the confidence level on the  $\alpha(t_0, \lambda, 880\text{nm})$  at different wavelengths. As  $\alpha(t_0, \lambda, 880\text{nm})$  depends on  $M_{OA}/b_{abs}(t_0, 880\text{nm})$  in Equation 13,  $\sigma_{\alpha(t_0, \lambda, 880\text{nm})}$  also does. We expressed  $M_{OA}/b_{abs}(t_0, 880\text{nm})$  as  $f_{OA}$ , using  $\sigma_{ATN}$  to estimate EC mass from  $b_{ATN}(880\text{nm})$ . At short wavelengths and low OA fractions, the confidence level on  $\alpha$  is within 0.1. However, with increasing  $f_{OA}$ , and at longer wavelength the uncertainty in predicting  $\alpha$  increases.

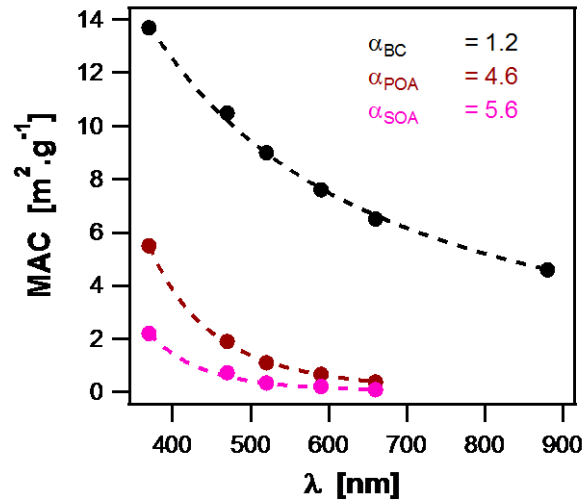
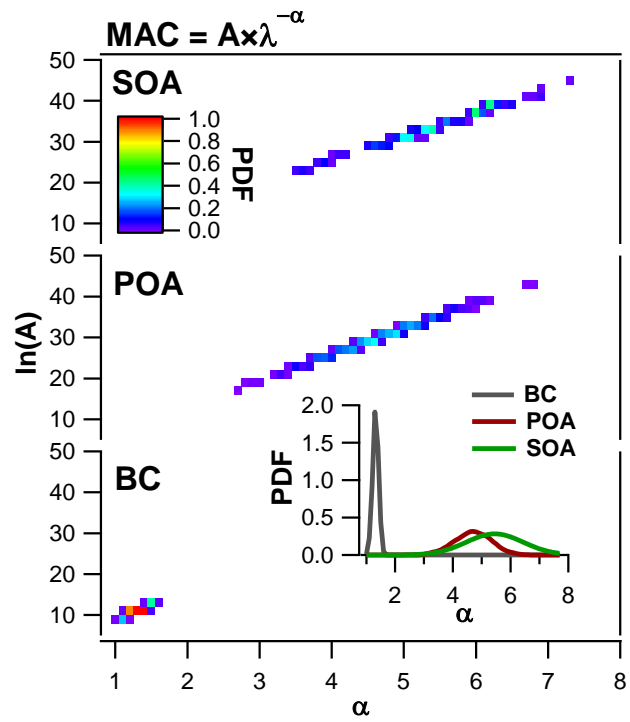
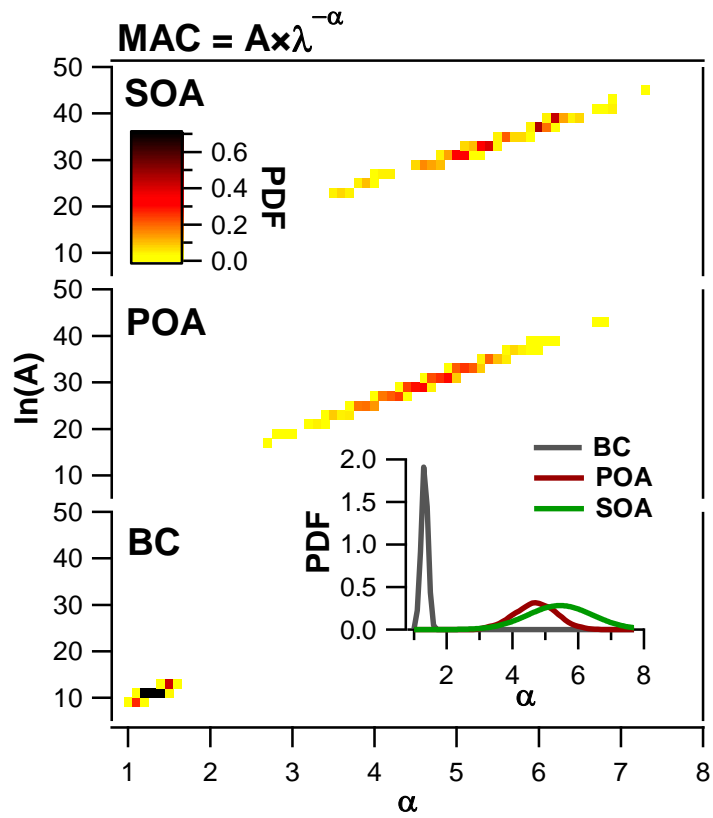


Figure S8: Power law fits through the average MAC of BC, POA and SOA calculated from Aethalometer measurements plotted as a function of wavelength. Note that  $MAC_{SOA}(880\text{nm})$  and  $MAC_{POA}(880\text{nm})$  are zero by definition.

Old Figure:



New Figure:



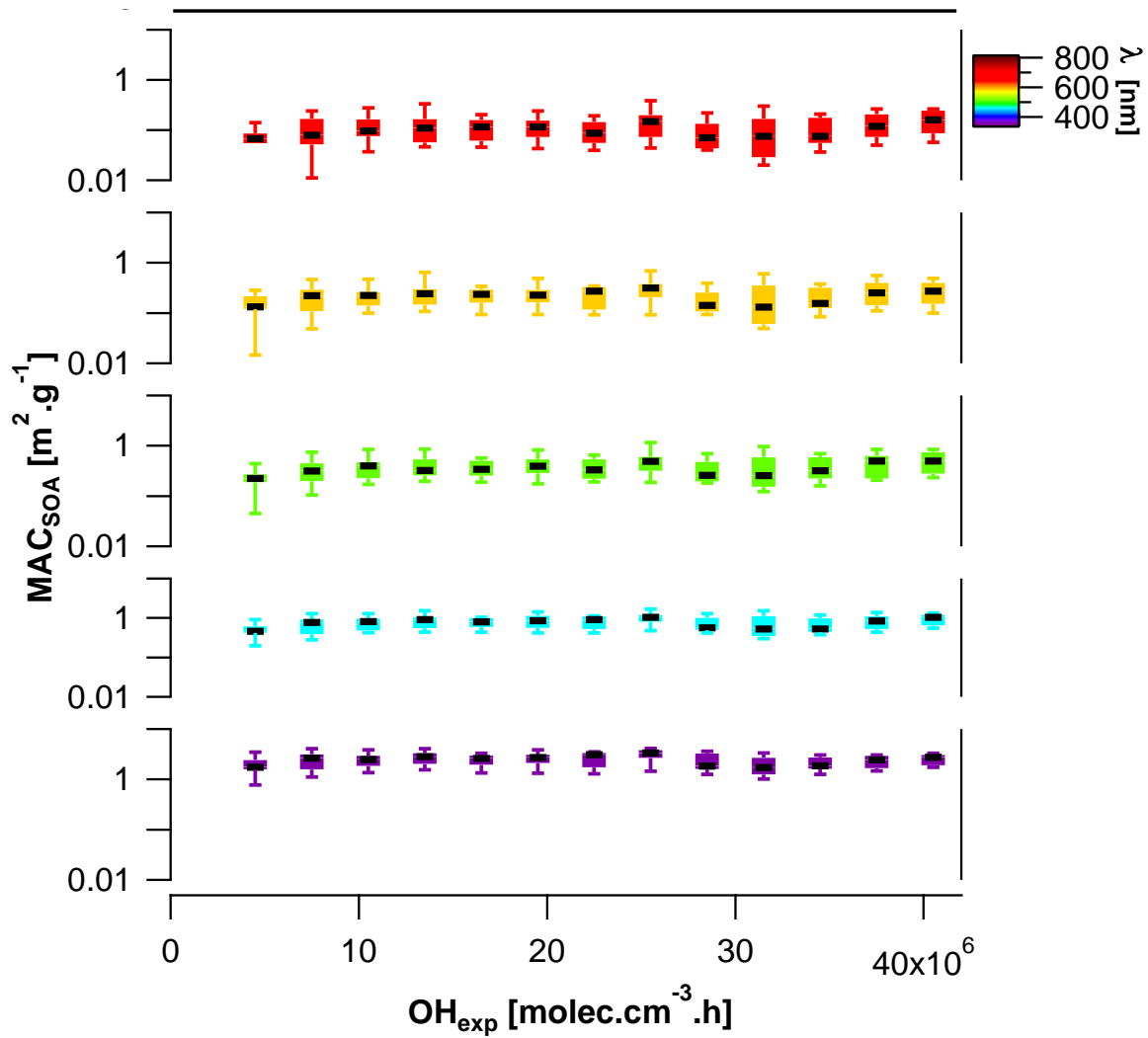
**Figure S9: Probability distributions of  $\alpha$  and  $\ln(A)$  describing the optical properties of BC, POA and SOA. Parameters for representing these distributions as a bivariate normal joint density function are shown in Table S1.**

The equation needed to generate these probabilities is  $f(\mathbf{X}) = \frac{|\Sigma|^{-\frac{1}{2}}}{2\pi} \exp\left(-\frac{1}{2}(\mathbf{X} - \boldsymbol{\mu})^T \Sigma^{-1}(\mathbf{X} - \boldsymbol{\mu})\right)$ , where  $\boldsymbol{\mu} = \begin{pmatrix} \mu_\alpha \\ \mu_{\ln(A)} \end{pmatrix}$ ; represents the average  $\alpha$  and  $\ln(A)$  values and  $\Sigma = \begin{pmatrix} \sigma_\alpha^2 & \rho\sigma_\alpha\sigma_{\ln(A)} \\ \rho\sigma_\alpha\sigma_{\ln(A)} & \sigma_{\ln(A)}^2 \end{pmatrix}$  is the covariance matrix. As  $\alpha$  and  $\ln(A)$  are determined from fitting the MAC vs.  $\lambda$ , their covariance is high. Therefore the selection of these parameters to represent the MAC profiles of BC, POA and SOA, should not be done independently but by using the probability density function above and the parameters in Table S1.

**Table S1: Parameters for normal joint density function.**

	BC	POA	SOA
$\boldsymbol{\mu}$	$\begin{pmatrix} 1.2 \\ 10.0 \end{pmatrix}$	$\begin{pmatrix} 4.6 \\ 29.2 \end{pmatrix}$	$\begin{pmatrix} 5.6 \\ 33.3 \end{pmatrix}$
$\Sigma$	$\begin{pmatrix} 0.09 & 0.054 \\ 0.054 & 0.61 \end{pmatrix}$	$\begin{pmatrix} 0.64 & 2.6 \\ 2.6 & 4.1 \end{pmatrix}$	$\begin{pmatrix} 1.36 & 11.4 \\ 11.4 & 8.35 \end{pmatrix}$

Old Figure:



New Figure:

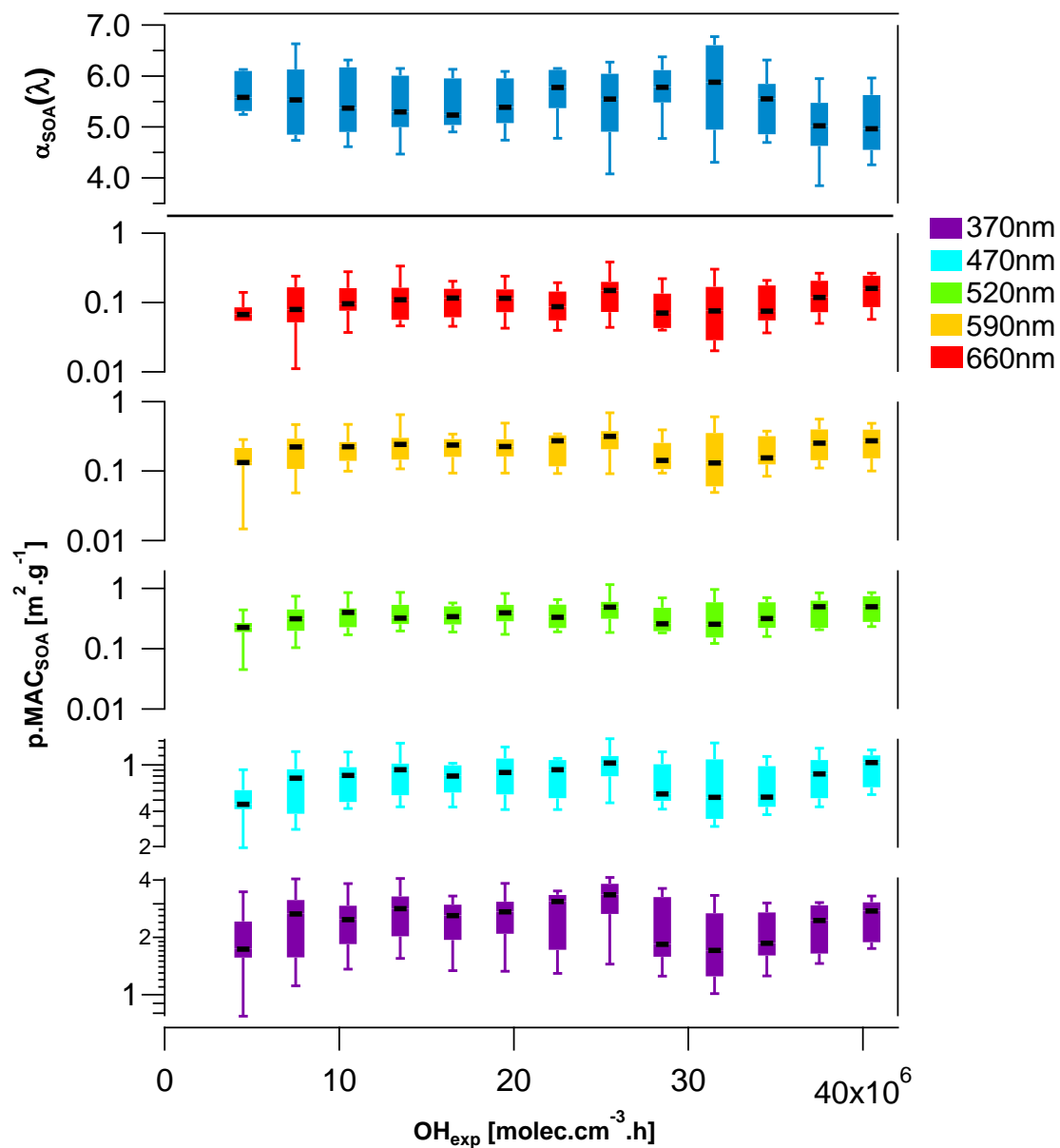


Figure S10:  $MAC_{\text{SOA}}$  as a function of OH exposure color coded according to the wavelength.

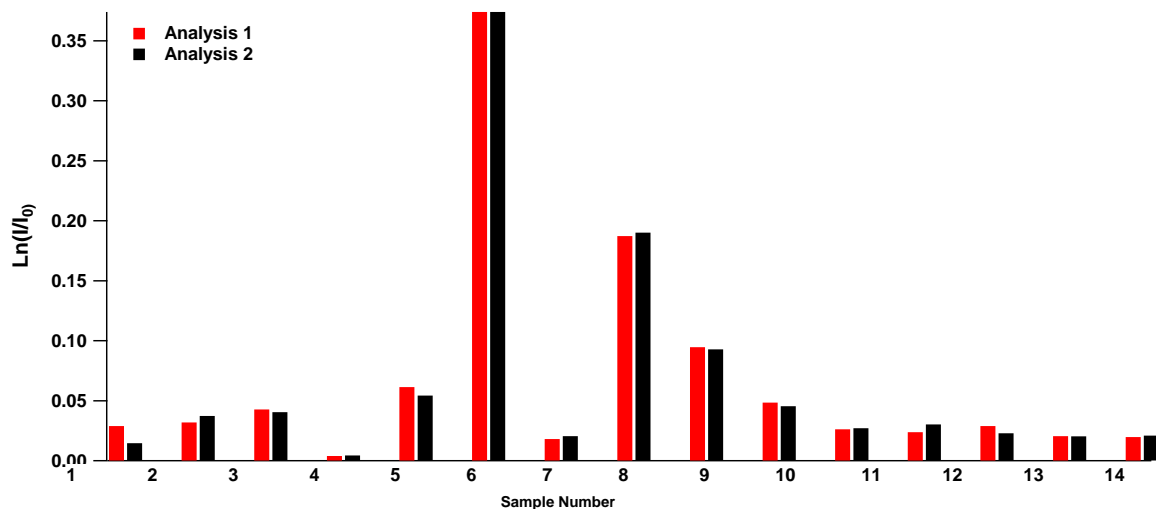


Figure S11: Absorbance measurements from UV-visible analysis of water extracted filters from several wood burning experiments showing very good repeatability (consistent within 10%).

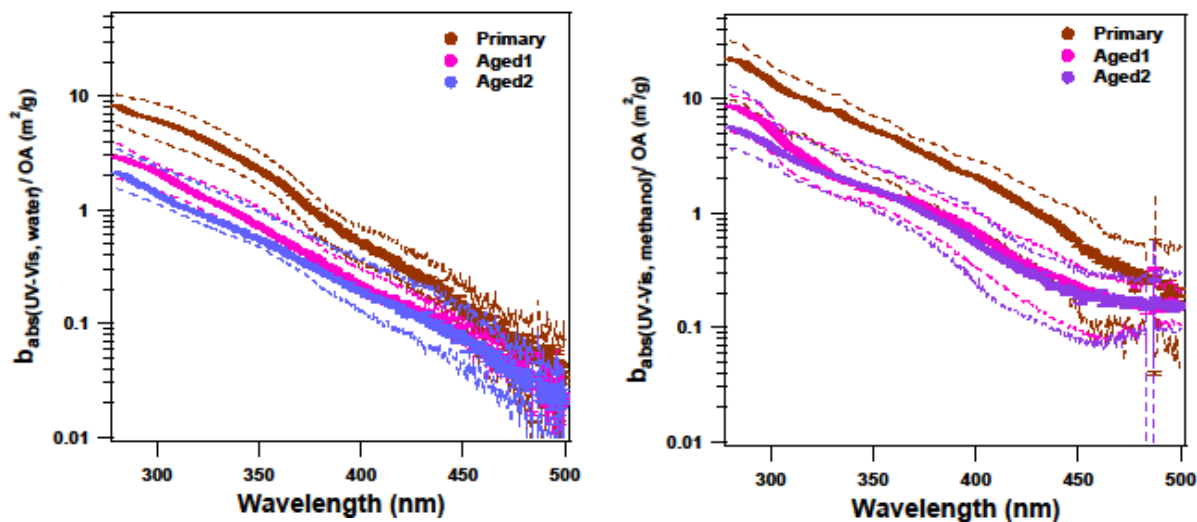
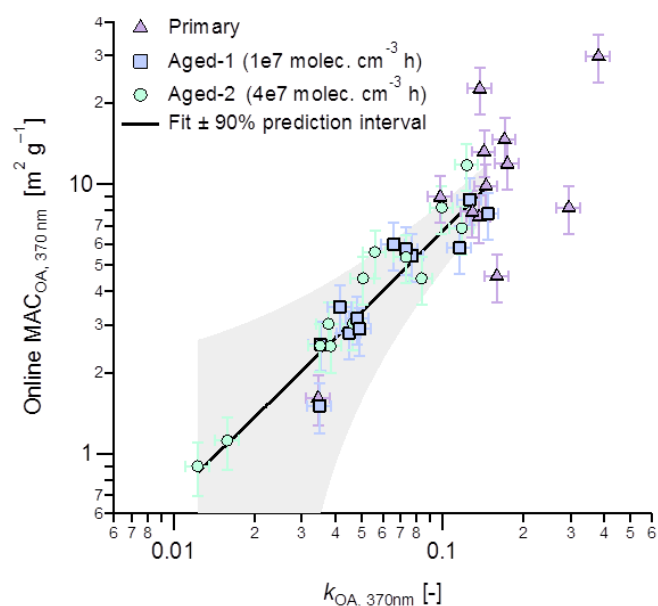


Figure S12:  $MAC_{\text{bulk}}$  (bulk absorbance of extracts normalized to AMS-measured OA) of primary, slightly aged (Aged1, OH exposure  $\sim 1 \times 10^7$  molecules  $\text{cm}^{-3}$  h) and aged emissions (Aged2, OH exposure  $\sim 4 \times 10^7$  molecules  $\text{cm}^{-3}$  h) for (A) water and (B) methanol extracts. The bold lines indicate the medians, and the dashed lines mark the 25<sup>th</sup> and 75<sup>th</sup> percentiles. At 450–500 nm, the methanol-extract absorption shows a constant absorptivity feature which was not present before aging, suggesting that the absorbing species may be partially-oxidized (partially-solubilized) primary OA, or reflect light-absorbing SOA.

Old Figure:



New Figure:

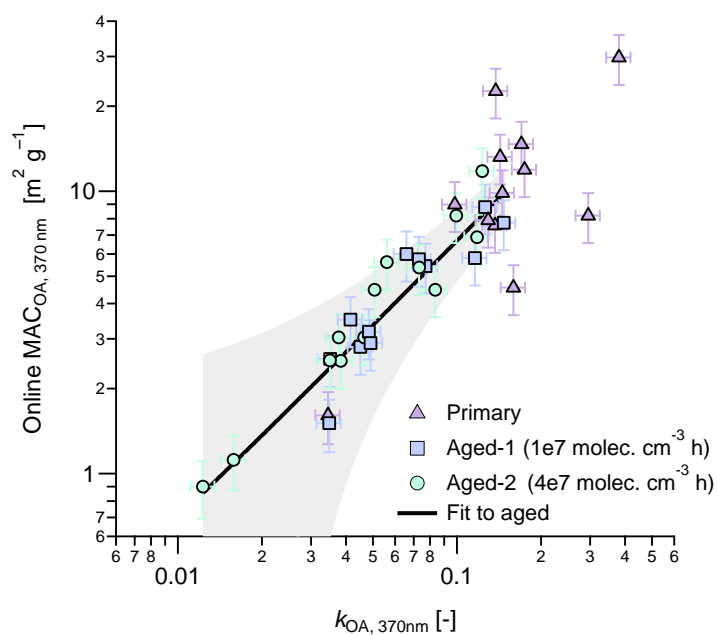


Figure S13: MAC<sub>OA</sub> at  $\lambda = 370$  nm calculated from aethalometer measurements vs.  $k_{OA}$  at  $\lambda = 370$  nm from the UV-visible measurements of the methanol extracts.

Figure S13: MAC<sub>OA</sub> at  $\lambda = 370$  nm calculated from aethalometer measurements vs.  $k_{OA}$  at  $\lambda = 370$  nm from the UV/visible measurements of the methanol extracts. The shaded region shows the 90% confidence interval of a weighted orthogonal regression (slope  $66 \pm 9$  m<sup>2</sup>g<sup>-1</sup>, intercept  $0.0 \pm 0.3$  m<sup>2</sup>g<sup>-1</sup>) to illustrate the relatively small range of variability in the data for aged samples.

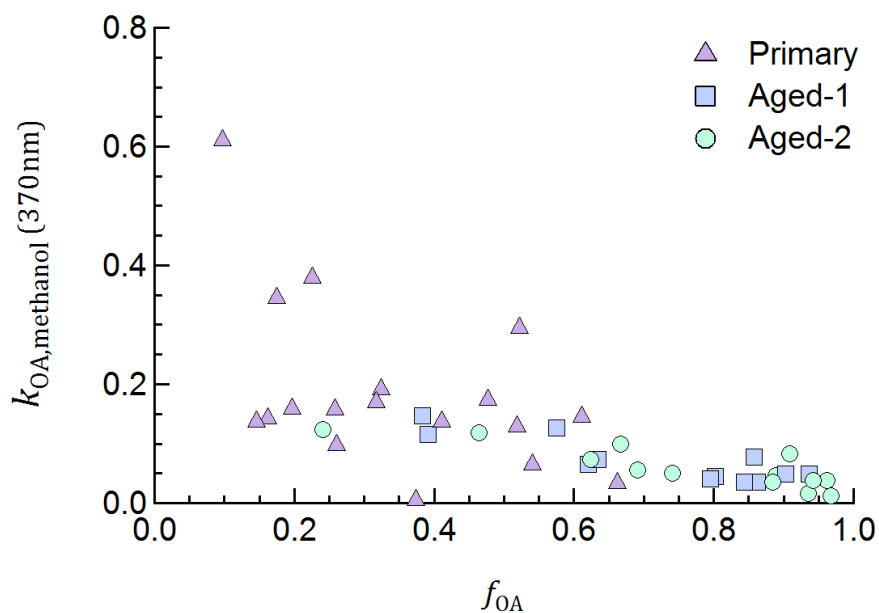


Figure S14. Similar to Fig. 7 in the main text, but plotted against  $f_{OA}$  for comparison to the other figures in this work.

## References

Pierce, J. R., Engelhart, G. J., Hildebrandt, L., Weitkamp, E. A., Pathak, R. K., Donahue, N. M., Robinson, A. L., Adams, P. J. and Pandis, S. N.: Constraining particle evolution from wall losses, coagulation, and condensation-evaporation in smog-chamber experiments: optimal estimation based on size distribution measurements, *Aerosol Sci. Technol.*, 42, 1001–1015, doi:10.1080/02786820802389251, 2008.

Bond, T. C., Habib, G. and Bergstrom, R. W.: Limitations in the enhancement of visible light absorption due to mixing state, *J. Geophys. Res. Atmos.*, 111(20), 1–13, doi:10.1029/2006JD007315, 2006.