

REPLIES TO REFEREES

We thank the referees for their insightful comments and suggestions that have helped us improve our manuscript.

We have answered to each of the referee's comments below. The reviewers' comments are shown in **bold**, and the text that has been added to, or modified in, the revised manuscript is shown in *italics*. The changes in the revised manuscript are shown using the Word 'Track changes' feature. The page and line numbers given in the answers refer to those in the ACPD version of the manuscript.

Reply to Referee #1

Summary: This paper presents long-term measurements (Nov. 2012 – Feb. 2015) of particle number-size distributions (PNSDs) from a rural background site in western Saudi Arabia. Authors used a twin DMPS to measure PNSDs in the size range from 7 to 850 nm and Vaisala WXT sensors to measure meteorological parameters relevant to new particle formation (NPF). Atmospheric NPF, growth, and shrinkage are not new findings and have been reported by several other investigators around the globe (e.g. Young et al. Yao et al., and so on...). Authors have discussed almost all these studies. But, this paper offers an important addition to the global aerosol dataset from a site, which has not been studied in the past and therefore, such long-term studies of aerosol measurements should be encouraged. I recommend this paper for publication in ACP only after authors have satisfactorily addressed both major and minor concerns below.

General comments: Page 3 and 11: Authors discuss by and large two mechanisms for aerosol shrinkage. First, the evaporation of semi-volatile organic vapours under favourable environmental condition. Second, transported smaller size particles to the measurement site. But authors fail to demonstrate either of the mechanism and said that “cause of the decreasing mode diameter (DMD) events observed in Hada Al Sham is further investigated in a future study”. Authors came up with the new term “decreasing mode diameter” to explain aerosol shrinkage but did not reveal the cause. I strongly suggest NOT floating new terms in the literature without clearly demonstrating it. In fact, authors could look at coagulation sink, if the smaller particles are really being transported to measurement site (obviously not too far away from the site, is there any primary source of these small particles in the vicinity of the site?). If the smaller particles are being advected to the site during shrinkage then I would probably expect high coagulation sink during shrinkage than growth. The simple ratio of Aitken to accumulation particles during growth and shrinkage event may be useful to demonstrate it. Further Authors could also calculate shrinkage rate, similar to growth rate, this would help future investigators for comparison.

We use the term decreasing mode diameter exactly because of the reason that we are not able to demonstrate the cause. In the manuscript (P3L3), we state that in our opinion ‘aerosol shrinkage quite directly implies a reduction in the size of individual particles’, but at the same time we feel that neither the previous analyses on this subject, nor the one presented in this manuscript, are sufficient to demonstrate that shrinkage would indeed be the cause. Therefore, we feel that the term ‘shrinkage’ has been used without clearly demonstrating it, and in order to avoid this, we try to refer to the phenomenon as objectively as we can (i.e. without implying the cause). There should be no denying that during these events the mode diameter is indeed decreasing. The intent of referring to this phenomenon as ‘DMD’ is therefore not to come up with a new term that we expect others to adapt, but only to use an expression that is suitable for the cause.

In this manuscript, our aim is to report on the general characteristics of NPF at the site. We felt that this had to include some discussion about the DMD events, as they are such a common feature of the NPF events here, even though our aim is not to explicitly examine the causes of DMD. We therefore wish to detain from further analysis on the subject here, while we do plan on examining the DMD events in an upcoming manuscript that specifically focuses on this subject.

Page 7, Line 26: Authors state that “undefined days showed some features representative of NPF events but these features were not clear enough for the days to be classified as NPF days”. I suggest referring to Buenrostro Mazon, S., et al., (2009) for classifying undefined events and classify undefined days according to Buenrostro Mazon, S., et al. approach.

We are aware of the guidelines presented by Buenrostro Mazon et al., (2009) to classify undefined days. Regardless, we decided to focus on the extreme ends of the ‘NPF spectrum’ mainly because of two reasons: (1) This simplifies the analysis and makes the output of the research more clear and straightforward to the reader. Especially since we already have an addition to the traditionally used event – non-event separation due to the DMD phenomenon. (2) The NPF characteristics (growth rate, formation rate and onset times of different phases) which are the main focus of this study are reliably obtainable only for the well-defined NPF class. In addition to these two reasons, the classification by Buenrostro Mazon et al., (2009) requires measurements of SO₂ and NO_x for the identification of pollution peaks, and we do not have this data. We, however, noticed that the sentence referenced by the Referee actually contains no information as it basically just states the definition of undefined days, which has been already accomplished earlier in the manuscript. The purpose of the sentence was to state that it seems the majority of the undefined days are actually NPF days that just fail to achieve the traditional requirements (referred to as ‘failed events’ by Buenrostro Mazon et al., (2009)), instead of particle number concentration peaks caused by other reasons. The modified sentence now reads:

“It should be noted, that the majority of the undefined days were likely days when new particles were formed, but continuous growth of particles was not observed at Hada Al Sham due to unfavorable meteorological conditions (referred to as ‘failed events’ by Buenrostro Mazon et al. (2009)).”

DMD events are more frequent during the summer, possibly suggests the contribution of evaporation of semi-volatile compounds to the diameter shrinkage, but it needs to be investigated. E.g. Page 11, line 1-5.

This will be investigated further in the follow-up paper about the DMD events. We did add a sentence connecting these observations to the seasonal variation of DMD events (as also suggested by Referee#2). The new sentence on P11 L2 now reads:

“This could also explain the observed summer maximum in the DMD event frequency, since these variables, which are likely to promote particle evaporation, obtain their largest values during summer.”

Other comments: Page 1, Line 27: You meant to say “large number concentration of primary particles” not “large size primary particles”. Remove “large”.

The original wording was actually intentional and its intent was to emphasize the contribution of large particles (say dp>100nm) towards the condensation sink, but since we do not explicitly show the contributions of different sized particles to the CS we changed the wording as follows:

“... likely reflecting the common anthropogenic sources of NPF precursor vapors and primary particles affecting the condensation sink.”

Page 1, Line 29: Authors state that “the NPF events in Hada Al Sham are exceptionally frequent and strong” but authors did not quantify how strong NPF events are? You may want to refer to Stanier et al., 2004 for classifying NPF events into strong, moderate and weak events based on the net increase of N25 during the first few hours of the event.

In this manuscript, we consider aerosol formation and growth rates as indicators of the strength of the NPF events. This is also explicitly mentioned in the text on P5 L10-13. On P11 L30 – P12 L2, we compare the annual median formation and growth rates to those obtained from long-term observations around the world and conclude that our values are very much in the high end of these observations, giving them a clear frame of reference. We do, however, agree that only stating NPF events to be strong is quite ambiguous and changed the wording on P1 L29 to specify our meaning:

“... the NPF events in Hada Al Sham are exceptionally frequent and strong both in terms of formation and growth rates.”

Page 2, Line 9: Please give the range for

Merikanto et al. (2009) report that 45% (uncertainty range: 31 – 49%) of global low-level (460 – 1100 m a.g.l) CCN at 0.2% supersaturation originate from nucleation, while Gordon et al. (2017) provide a corresponding estimate of 54% (38 – 66%). Yu and Luo (2009) do not report global averages nor uncertainty ranges but show zonally-averaged latitudinal and vertical distributions of annual mean values of the secondary fraction of CCN0.2% (supp. material) ranging mostly between 40 and 70% in the low-levels with the global average seemingly in the 50-60% range. Thus, we will give the uncertainty range based on the estimates by Merikanto et al. (2009) and Gordon et al. (2017).

“... they are estimated to produce around half (31–66 %) of the cloud condensation nuclei (CCN) in the lower troposphere...”

Page 2, Line 10: Several studies highlight the importance of NPF in heavily polluted megacities around the world, especially developing nations. Authors should consider citing them here. A few of them are; Stanier et al., 2004, Kanawade et al., 2014, Yue et al., 2009, Xiao et a., 2015, Iida et al., 2008, Yu et al., 2017 and so on...

Here we wanted to cite studies that have specifically evaluated the contribution of secondary vs primary particles in polluted regions. Although the studies suggested by the Referee do point towards the importance of NPF in polluted regions, we could not find this kind of quantitative evaluations included in them. Yet, we do feel that the overall occurrence of NPF in polluted cities could be mentioned explicitly (although it is implied in the previous sentence). In addition, we added a new reference that discusses the potential increase in CCN concentration due to NPF (Yue et al., 2011). The modified sentence now reads:

“Even in polluted regions, where the primary emissions are high, NPF events are found to occur frequently (Yu et al., 2017) and they are estimated to be a significant contributor to the particle number concentrations (Yue et al., 2011, Kulmala et al., 2016; Yao et al., 2018).”

Page 2, Line 30: Please could you specifically state the importance of long-term measurements.. e.g. to reduce uncertainties in aerosol nucleation rates which are currently a few orders of magnitude in global models, CCN estimation from NPF in the boundary layer or troposphere which also show large range.

Added more specific information about the importance of long-term PNSD measurements:

Long-term measurements are needed for obtaining reliable estimates on the average properties and seasonal tendencies of atmospheric NPF. Such data are essential in evaluating the performance of global models which currently have large uncertainties in simulating atmospheric NPF, as well as its contribution to CCN budgets and aerosol radiative effects in different environments (Pierce and Adams, 2009; Makkonen et al., 2012; Gordon et al., 2016; Semeniuk and Dastoor, 2018).

Page 3 Line 29-30: How far each emission source and in which direction from the measurement site? What is the elevation of the site above mean sea level?

Added more specific information about the emissions sources in the region (P3 L28-31):

“Jeddah, the second largest city in Saudi Arabia, is located by the Red Sea ~ 60 km to the west from Hada Al Sham. In Jeddah there are several major emission sources, including power generation plants, a sea water desalination plant, as well as an airport and a harbor, both of which experience heavy traffic due to the combined effect of economic growth and the vicinity of Makkah, located ~ 40 km to the south from Hada Al Sham. Other major emission sources in the region include a petrochemical refinery and steel industry in Rabigh (~ 130 km NW, see Fig. 3), and a large oil refinery in Yanbu (~ 290 km NW).”

Added information about site elevation (P3 L24):

“Hada Al Sham (21.802 °N, 39.729 °E, 254 m a.s.l) ...”

I would suggest replacing “NPF start” by “NPF event start” everywhere so that “NPF end” refers to the time when aerosol nucleation process ceases and not the entire NPF event. In that case, figure 2 legend should be “NPF+DMD”, “NPF event”, “Undefined event” and “Non-event” and elsewhere in the text and figures (fig. 6 and so on).

Our choice of terms in describing the event times was based on the idea that now ‘NPF start’ and ‘NPF end’ describe the starting and ending times of the same physical process which is the formation of new particles (in this case production of 7nm particles). By changing ‘NPF start’ to ‘NPF event start’, this connection would be partly lost. While we do agree that using ‘NPF event start’ would better tie together the start and end of the whole NPF process, we feel that this change would not significantly reduce the risk of a mix-up between ‘NPF end’ and ‘NPF event end’, which is the main problem here. Thus, we think that clearly indicating our use of terms both in the text as well as in the figure captions should be enough for the reader to understand our terminology.

Are aerosol number-size distributions measurements are corrected for diffusional losses in the sampling line or diffusion dryer? What are the dimensions of the diffusion dryer? I assume the smaller particle losses are not more than 5-10

Aerosol number-size distributions measurements were corrected for diffusional losses in the inlet line and they were around 15 % at lower size limit of the DMPS measurements. The dimensions of the diffusion dryer were 1 m × 10 mm. We added a sentence on P4 L16 mentioning the losses for the smallest particles:

“PNSD measurements were corrected for diffusional losses in the inlet line and they were at maximum around 15 % at the lower size limit of the DMPS.”

Figure 1: As a reader, I find difficult to follow this figure. Why do you cite Hussein et al., 2005 in the figure caption? If you have used GR calculation methodology based on Hussein et al., 2005 then please discuss it in the methods section (which you have!). What are “three” different black dots, one obviously mode diameter, the bottom one, what is middle and the top which is very sparse? It would also help the reader if you could use nanometer for y-axis. Figure

legend says black dots as mode fit, I think, its mode diameter; the black line is the mode fit. I also suggest describing vertical lines in the figure caption than showing in legend. I can not figure out “NPF event” end time visually, though can see “NPF” end time. It would help if you could over-plot let’s say $N < 25\text{nm}$ (there is hardly any primary source of particles less than this size unless we measure next to the source e.g. vehicle exhaust).

The citation discusses the automated fitting algorithm, which is also discussed in the methods section.

At each measurement time, the aerosol number size distribution can consist of several different modes which can often be described with reasonable accuracy using log-normal distributions. The typically found modes in the $dp < 1\mu\text{m}$ range are an Accumulation mode, Aitken mode and a Nucleation mode (which is seen especially during nucleation events). Because of this, the model is also allowed to use several (in this case 2–3) modes (i.e. log-normal fits to the PNSD) as it attempts to describe the PNSD (as described in the methods section P5 L17–19). The different black dots seen in the figure are the geometric mean diameters (GMD) of these fitted modes. Therefore, in this context the word ‘mode’ does not refer to the single diameter with the largest $dN/d\log Dp$ -value (as it would in a purely statistical sense), but rather to the whole population of particles described by a single unimodal log-normal distribution. We do agree with the Referee that describing the black dots as ‘mode fit points’ is inaccurate, and their description is now changed to ‘GMD of fitted mode’ in figure 1 and elsewhere in the text.

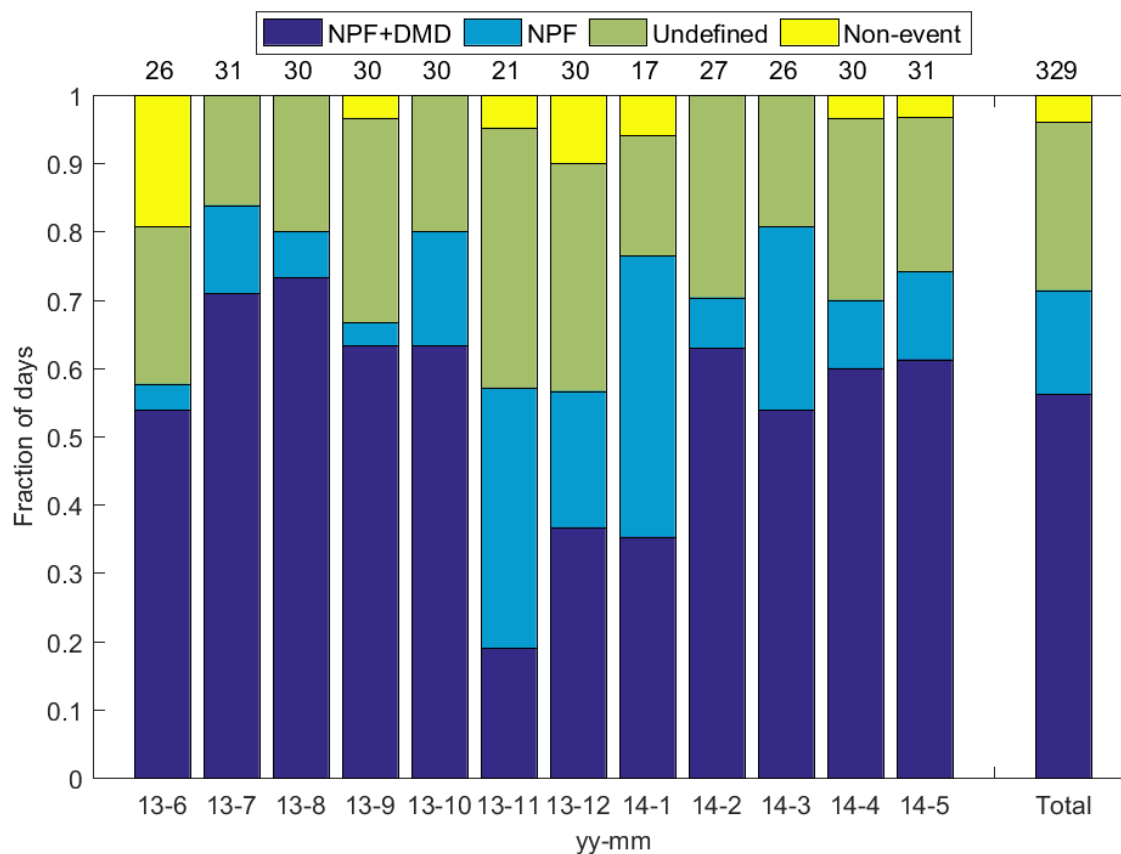
We agree that the NPF event end time is not easily distinguishable in this case (in several cases the event end time is much more clear as might be seen in Fig. A1) and we do not claim to have a strong argument for the exact placement of said time. In this specific case the NPF event seems to end somewhere between say 18:30 and 21:30 and beyond this level of accuracy the decision is quite subjective. We would like to avoid plotting additional data on the figure since it is already quite busy, as also implied by the Referee. In addition, this could be somewhat misleading since no number concentration data (other than the visual interpretation of the surface plot) were actually used for determining the times. In the scope of this manuscript, the main point of determining the NPF event end times is just to point out the fact that a large fraction of the NPF events do indeed end in a relatively short time from the NPF start, which implies a spatially limited NPF area (as discussed on P11 L6 onwards).

Based on the comments we’ve made the following changes:

- Removed reference to Hussein et al. 2015 from the figure caption
- Changed the y-axis to nanometers (also in Fig. A1)
- Replaced ‘mode fit (point)’ with ‘GMD of fitted mode’
- Added description of vertical lines in the figure caption (+added thin black lines on the edges of the colored vertical lines for clarity, as a response to a comment by Referee#2)

“The figure also illustrates the times describing the progression of the NPF events with colored vertical lines (NPF start – blue, NPF end – orange, DMD start – yellow and NPF event end – purple) and the geometric mean diameters (GMDs) of fitted modes with black circles. The GMDs selected for the calculation of the growth rate are shown using white circles, and the black curve shows the linear fit to these points.”

Page 7, Line 29: Authors state that “no clear seasonal pattern is observed in NPF frequency”- could it be that 5 months data is available only for one year whereas the remaining months data is available for 2 or more years (e.g. Feb. 70 days versus June 26 days). This can be ensured by selecting a year during which all months data is available. I would be interested to see figure 2 for the year during which all months data available.



In the figure above, we have calculated the monthly event type fractions for a continuous 1-year period from Jun 2013 to May 2014. Even by selecting this period, no clear seasonal pattern is observed in the NPF frequency. We compared each of the monthly NPF fractions with the total NPF fraction (chi-squared test comparing two proportions), and found out that it is *not highly unlikely* that the monthly observations would share a same NPF frequency as the total observations (lowest p-value being $p=0.091$ for Dec 2013). The same was also true for the observations shown in Fig.2 in the main text (lowest p-value being $p=0.055$ for Dec). Our interpretation of this result is that if none of the monthly NPF fractions is significantly different from the annual average fraction, it should be unlikely that any patterns in the monthly fractions would be meaningful.

Figure 4: Since authors sub-divided NPF event days into DMD and non-DMD events, I suggest to include air mass distribution maps for DMD and non-DMD events?

(See answer to first question):

In this manuscript our aim is to report on the general characteristics of NPF at the site. We felt that this had to include the division into DMD and non-DMD events even though our aim here is not to explicitly examine the causes of DMD. We therefore wish to detain from further analysis on the subject here, while we do plan on examining the DMD events in an upcoming manuscript that specifically focuses on this subject.

Authors report growth rates in the diameter range of 7 to 12 nm (7.4 nm h^{-1}). How many bins do you have in this size range? From the figure, I can see that the fit line reaching upto 40 nm or may be more so that the reported GR are not actually GR of particles in the size range from 7-12 nm? Please clarify or correct.

This issue is specifically addressed in the manuscript on P5 L24–28: “The growth rates presented in this study were determined so that they would best describe the growth of particles in the diameter range 7–12 nm, as this is the range used in the determination of the formation rates. In principle, this could be done by always selecting only the mode fit points below 12 nm, but since single points are subject to fluctuations in the PNSD and the number of fitted points below the 12 nm threshold was often quite small, points above 12 nm were also often included into the fit to obtain a more robust estimate of the particle growth.”

We added a sentence on P5L28 to clearly state when the inclusion of additional points was concerned:

“This was, however, done only if the growth rate seemed to remain constant above the 12 nm threshold.”

The number of bins below the 12 nm threshold is 7. An additional reason for selecting fit-points outside the 7–12 nm range, that is currently not mentioned in the text, is the occasional poor performance of the automatic fitting algorithm (as can be seen in the example figure (Fig. 1) at t=11:00, dp=20nm, where the model seems to fail in capturing the development of the growing mode).

I would suggest over-plotting particle mode diameter (thin black line) in Figure A1. There are some data gaps as you mentioned in the text which is not visible on this figure.

There are actually very few gaps in the data during these selected periods. We changed the figure so that the data gaps are shown in white. In addition we changed the y-axis to show nanometers instead of meters as suggested earlier by the Referee.

I have not checked the paper for linguistic/typo errors so I suggest authors to take off those carefully if any.

The manuscript was checked for linguistic/typo errors. At least the following corrections were made:

- Unified the spelling of ‘number size distribution’ (i.e. replaced appearances of ‘number-size distribution’ and ‘number size-distribution’)
- Unified the spelling of ‘Saudi Arabia’ (replaced ‘Saudi-Arabia’) on P1 L19
- P2 L25, replaced ‘In addition’ with ‘Interestingly’ to avoid reoccurring use of ‘in addition’ and to highlight the exceptionality of this observation.
- Other minor changes considering choice of words and the excessive use of ‘the’ (reviewable via track changes)

Reply to Referee #2

The paper by Hakala et al. (2019) describes the occurrence of new particle formation (NPF) at a rural background site in Saudi Arabia, Hada Al Sham, using a two-year long dataset. This study is of high interest, as it reports observations from a still poorly documented region / environment, where anthropogenic emissions are likely to play a significant role in atmospheric processes. More broadly, such long timeseries are needed for a better understanding of NPF, and in turn better description of the related effects on climate in global models. Moreover, the paper is well written, very pleasant to read, and figures are clear. I would however place a caveat on this analysis, since the investigation of a specific aspect of the observed events (DMD) strongly contributes to the interest / novelty of this work, and is unfortunately not complete to leave room for a companion study. Nonetheless, I recommend the publication of this study after some minor revisions which are listed below; they concern the main text, but abstract and conclusion should be modified accordingly.

P8, L11-12, L16: “NPF event frequency has been shown to be affected by at least: solar radiation, SO₂ concentration...”, “which is widely regarded as the driving compound of atmospheric new particle formation”. I would suggest to slightly balance these statements which are too strong in my opinion, since SO₂ (and in turn H₂SO₄) has not been shown to be a limiting/determinant precursor in all environments, as well stressed on P2, L23-26.

We changed the wording of our statement and added a few references to cover a wider range of studies and environments.

“... ,which is regarded as the driving compound of atmospheric new particle formation in most environments (Weber et al., 1997;Birmili et al., 2003;Kuang et al., 2008;Paasonen et al., 2010;Yao et al., 2018)”

P9, L16-17: The reported results strongly point toward a significant / dominant role of anthropogenic precursors, but, again, this assessment (“implies that”) is in my opinion too strong. Indeed, I think that a positive influence of marine conditions on NPF, even if minor, cannot be excluded based on the available measurements, since the Red Sea sector / coastal area is also a signature of the air mass back trajectories on event days. One may for instance hypothesize that marine conditions could affect NPF:

- 1) with some specific precursors;
- 2) but also because they might present favourable conditions for NPF to be triggered, such as for instance lower CS compared to “pure continental” air masses;
- 3) or because they most likely display increased RH compared to inland air masses, which might contribute to higher NPF frequencies / particle formation and growth rates observed in western air masses, as discussed later in the paper.

We agree that the word ‘imply’ conveys too much causality, especially at this point of the manuscript, and changed the wording as follows:

“The fact that all of the non-event days are observed when the air masses are coming from the inland suggests that the NPF events observed in Hada Al Sham might be related to the emissions from the coastal anthropogenic activities. In addition, the influence of marine air could be beneficial for NPF occurrence e.g. due to lower condensation sink or some specific precursors.”

We do however argue that our results later on clearly point towards the importance of the anthropogenic emissions over the marine conditions since:

- 1) Night time and early mornings are very calm with slightly easterly winds -> It is unlikely that marine conditions or precursors would play a role here (distance to the coast is 60-70km), yet NPF does start
- 2) CS is higher on NPF days than on non-NPF days (fig. 5). Also, FR and GR increase with CS, which suggests positive contribution from pollution sources rather than from the clean marine environment.
- 3) NPF seems to be spatially limited in the western direction -> no NPF at all is observed in ‘marine-enough’ air masses

All of this does not mean that marine air mixed with the anthropogenic emissions could not be the *optimal solution* (e.g. due to the RH effect), but the anthropogenic emissions definitely seem to be the crux of this equation. Because of this, we would like to detain from emphasizing the potential benefits of marine air masses to NPF in the conclusions and the abstract. Regardless, we did modify a sentence in the abstract (P1 L22:” In Hada Al Sham, the precursor vapors seem to be related to the transport of anthropogenic emissions from the coastal urban and industrial areas, since no NPF events are observed in air masses coming from the sparsely inhabited inland”) that contained the same implication as the one pointed out by the referee. The modified sentence now reads:

“Several factors suggest that in Hada Al Sham these precursor vapors are related to the transport of anthropogenic emissions from the coastal urban and industrial areas.”

We also reformulated a sentence on P10 L17-18 regarding these changes by changing ‘depend on’ to ‘are related to’:

“This is in agreement with the interpretation that the NPF events observed in Hada Al Sham are related to the transport of emissions from the coastal regions.”

P9, L31-32: Could the authors comment more on the results they report for CS? In specific, what could be the reason for higher nocturnal values on event days? Could this observation suggest an enhanced accumulation of the precursors during the same nights, thus facilitating the occurrence of the process on the next morning? This would be consistent with the fact that the sources and sinks of the driving precursors share the same origin, as suggested on P13.

Yes, this is also our interpretation for the higher CS values during the nights and early mornings of NPF days. As mentioned by the referee, we do make this connection on P13 L27-30, but we do agree that this speculation could be initiated already when discussing Fig. 5. Since the paragraph starting on P9 L26 focuses on discussing the effect of PM10, we added a new paragraph after this discussing CS. We also included a suggestion made by the referee in the later comments about drawing a parallel between a mountain site, where similar observation is made.

“Higher CS values during NPF days (Fig. 5c) are also reported from a high altitude site in the Swiss Alps (Boulon et al., 2010). Here this connection is speculated to stem from the coupled appearance of NPF precursor vapors and CS due to their common lower altitude sources. Similar situation could apply to Hada Al Sham in case both the CS and the NPF precursor vapors originate mainly from the same sources. The high CS during the calm nights and early mornings of NPF days could then suggest enhanced accumulation of precursor vapors, thus facilitating the occurrence of the NPF process after sunrise.”

We also changed the wording on P9 L31 about ‘CS not being a determining factor’ to ‘CS not being an inhibiting factor’, since in some sense, albeit indirectly, CS can be seen as a determining factor.

L34: “did not have a significant effect on these results”: Can the authors give an average of the CS increase observed when including APS measurements in the calculation?

The median increase in CS when including APS measurements was 3.7% with the 10th and 90th percentiles being 1.7% and 8.9%, respectively. The 90th percentile increase is similar to the increase obtained for the high PM non-NPF days, while the median increase is similar to that on NPF days. We modified the sentence on P9 L34 to give the approximate increase on non-event days, which is most important for the presented result:

“... did not have a significant effect on these results (increase in CS on non-event days was around 10 %).”

P10, L2-10: I would not restrict the conclusions to the PM10 related observations, and more broadly suggest the lack of precursors, not only anthropogenic, in the inland, in particular because it is clearly mentioned later (P13, L19-21) that the enhancing effect of mineral dust has been previously reported in conditions (timescale) which differ from that of the present study.

We do not fully understand the meaning of this referee comment. The logic in our discussion here is that we first consider PM10 as sink for the NPF precursors, but then show that even if the precursor emission strength was similar on both NPF and non-NPF days, the sink should not explain this difference since the sink is higher on NPF days (even if we consider the effect of the PM10 to the CS). Therefore, the conclusion from this would be that the precursor emission strength was not similar for both cases, but significantly smaller for the non-NPF case. After this, we point out that there is even a possibility that high PM10 could enhance NPF, and since we do not see NPF, we are most likely missing precursor sources. The conclusion is therefore that we are missing precursor sources regardless of the effect of PM10 on NPF. We can't however determine the effect of PM10 on solar radiation and added a mention about this on P10 L4:

“In Hada Al Sham, the concurrence of high PM10 values and non-event days would then mainly highlight the lack of anthropogenic emissions in the inland, although we cannot quantify the effect of PM10 on radiation.”

However, the lack of precursors seems more likely in our opinion.

P10, L33 - P11, L1-4: The different characteristics of the DMD events are discussed throughout the paper, and I think that the reader would sometimes benefit from some clear links between the observations. For instance, the seasonal variation of the DMD frequency is reported on P8, L1-2, but is not further commented in this section. The analysis of the environmental conditions together with the timing of the events provided in the next section points toward an effect of temperature on the occurrence of DMD. This observation should afterward be used to further discuss the seasonal variation of the DMD frequency, which supports such an effect of temperature, since the maximum of the DMD frequency in summer coincides with highest temperatures.

Same type of comment also applies to the CS (P9, L31-32; P13, L24-30).

Added a sentence on P11 L2 making a connection to the seasonal variation of DMD events:

“Since all of these variables obtain their highest values during summer, these effects could also explain the seasonal variation in the DMD event frequency.”

Based on a previous comment by the referee, a clearer link was already established between the sections about CS on P9, L31-32 and P13, L24-30.

P12, L4-7: J and GR show a seasonal cycle in Hada Al Sham; in contrast, the NPF frequency does not, which is not “common”, as, for instance, Nieminen et al. (2018), report a seasonal cycle of the NPF frequency with a maximum during local spring / summer for 30 stations out of 36. This observation suggests that in Hada Al Sham, the occurrence and strength of the NPF process are somewhat disconnected, or not driven by the same “factors”. The fact that J and GR have the same seasonal cycle is also interesting, and, again contrasts with the results recently reported by Nieminen and co-authors. Indeed, they report instead similar cycles for the NPF frequency and J, while GR usually displays slightly different variations, which are, at least to a certain extent, attributed to the involvement of different vapours in the successive stages of the NPF process.

Could the authors comment on these aspects?

We agree that these seasonal cycles are different than in most locations that are reviewed by Nieminen et al (2018). We interpret the lack of seasonal cycle in NPF event frequency being caused by favourable conditions for NPF throughout the year, except during the strong easterly winds, which seem not have a seasonal preference. It seems that in Hada Al Sham the seasonal cycles of the variables favorable for NPF – solar radiation level and concentrations of SO2 and other precursors of (extremely) low volatile vapors – are not strong enough to decrease NPF frequency in any season (as briefly mentioned in the manuscript on P7 L2). Additionally, as SO2 (favouring NPF) and CS (expected to inhibit NPF) are both presumably dominated by similar sources, or at least sources in same areas, the ratio of particle sources and sinks may not vary through the year enough to cause observable cycle in NPF frequency

We assume the similarity between seasonal cycles in J and GR to be caused by similar sources of the precursors for extremely low volatility vapors participating in NPF (SO2 and presumably some bases and/or VOCs) and low volatility vapors responsible for the growth (presumably mainly VOCs). In Hada Al Sham we expect the role of biogenic emissions to be minor and thus both NPF and growth related vapor concentrations to be dominated by anthropogenic sources in the same areas, which makes it logical that J and GR would have the same seasonal cycles. In most parts of the world, J is likely

mostly controlled by inorganic (anthropogenic) emissions (SO₂) + radiation, while growth can be controlled by availability of biogenic VOCs. Thus, in these environments during spring the combination of e.g. heating emissions + low BLH + low temperature + ever increasing radiation could create the optimal conditions for J, while the maximum GR values would be observed only later in summer when the (mostly temperature-dependent) biogenic VOC emissions are largest.

To include some of this discussion in the manuscript, we modified the paragraph starting from P12 L4 and it now reads:

“Both the formation and growth rates show a similar seasonal cycle, with largest values during the summer and early autumn (Fig. 8a, b). For the growth rates, a summer maximum is often observed also globally, while the formation rates peak in most locations during spring (Nieminen et al., 2018). The different seasonal cycles of J and GR at these sites could result from different species controlling the initial formation and further growth of the particles. In many places, the summer maximum in growth rates might be related to increased photochemical activity and increased emissions of biogenic VOCs as a function of temperature (Yli-Juuti et al., 2011) while the formation of particles could be more dependent on anthropogenic emissions (Nieminen et al., 2014). Conversely, the similar cycles of both J and GR in Hada Al Sham could then suggest that here the precursor sources of particle-forming and -growing vapors are similar. Since the emissions from biogenic sources are expected to be minor, due to the lack of vegetation in the area, these sources are likely to be anthropogenic. In Fig. 8a we show the monthly crude oil consumption in Saudi Arabia which seems to peak around the same time as J and GR. Thus the increased emissions from energy production, resulting from the growing need for air conditioning during summer, could possibly explain the observed seasonality in J and GR, although increased solar radiation is also likely to play a role.”

P12, L28-33: Is RH on average lower on non-event days? I would expect so given the inland origin of the air masses on these specific days, but RH (and related effects on NPF occurrence) is surprisingly not discussed in Sections 3.1 and/or 3.2, despite being shown on Fig. 6.

RH is indeed on average lower on non-event days than on NPF days, but similar to NPF days in June, which is when a large fraction of the non-events occur. Even though the lower RH could contribute to the absence of NPF, as suggested by Fig. 9b, Fig. 9b also shows that NPF can occur even when RH is extremely low (RH<10%). Therefore, we do not consider RH as a limiting factor for NPF, and only discuss its effects in the context of Fig. 9 on P12 L28 onwards and on P13 L10-12.

P13, L7-10: I am not sure if the correlation between J and ABL height and, more specifically, the fact that events are observed earlier with respect to ABL development during summer time is related to the ABL height itself only. Based on Fig. 8.a, emissions from energy production are increased during summer. Assuming that these emissions directly affect the amount of vapours relevant to NPF, we may thus assume that there is a larger pool of these precursors available already before sunrise in summer, and that in turn NPF is mainly limited by photochemistry. This would be consistent with events triggered shortly after sunrise, and consequently also earlier during ABL development. During other seasons, NPF might in contrast be more vapour-limited, and thus started later, both with respect to sunrise and ABL development, when there is a sufficient amount of precursors. In addition, would it be reasonable to assume that during summertime radiation is stronger already in early morning, thus leading to more “efficient” photochemistry also contributing to earlier occurrence of the process?

In our text on P13 L7-10, we only point out that the lowest event-time ABL heights seem to be occurring during summer without providing any explanation (or cause-and-effect relationship) for this observation. The speculation presented by the referee does, however, seem like a reasonable explanation (at least the first part), and we decided to include it in the text on P13 L10:

“This observation could be caused by the higher emissions during summer (see Fig. 8a), since if the concentrations of NPF precursors are higher, the onset of NPF events is likely to be more sensitive to an increase in solar radiation.”

The second part, i.e. that the radiation would increase faster after the sunrise during summer, does not seem to be true. According to a simple radiation model, the early-morning increase rate in radiation peaks during spring and autumn, while it is very similar during summer and winter.

P13, L24-30: These observations are very similar to those reported from several high-altitude stations, where NPF is thought to be triggered from precursors originating from lower altitude and transported at the sites together with their sink. The authors could actually draw a parallel with this situation (eg: Manninen et al. 2010; Boulon et al., 2010). The fact that the sources and sinks of the precursors share the same origin also most likely explains (at least to some extent) why the CS is on average lower on non-event days compared to event days (P3, L31-32).

This is now discussed in the new paragraph added on P10 L6.

Technical / minor comments:

Title: Even if it was convenient to keep the title short, I have been afterward a bit surprised that the word “shrinkage” is used in the title, as the authors clearly explain on page P3 L2-4 why they decided to “avoid” it in the paper!

We do agree that using the word ‘shrinkage’ in the title is somewhat questionable, or at least incoherent with the text otherwise. Yet, it allows for much more simplicity in the title, as mentioned by the Referee, and it also ensures that the readers interested in the ‘aerosol shrinkage phenomenon’ will easily find and recognize this manuscript’s contribution to the topic. Because of these reasons, we would like to keep the title as is.

P2, L21-22: “These species are likely ... or anthropogenic VOCs”: could the author reformulate this sentence for clarity?

We reformulated the sentence and added one relevant reference (Dall’Osto et al., 2018). The modified sentence now reads:
“These species are very likely to be low- or semi-volatile organic compounds, formed by the oxidation of either biogenic or anthropogenic VOCs (Smith et al., 2008; Tröstl et al., 2016; Dall’Osto et al., 2018).”

Also, we changed the spelling of ‘X volatile’ to ‘X-volatile’ elsewhere in the text

P3, L12 and P4, L4: the dates reported for the start/end of the campaign are slightly different.

The first mention was describing the availability of DMPS-data, while the second mention reported the start/end of the whole campaign (although the ending months were still inconsistent). We fixed both to give the range for the data used in this study which is the duration of available DMPS-data: Feb 2013–Feb 2015.

P5, L25: “principle” instead of “principal”.

Fixed.

P6, L7: “due to collision and coalescence”.

Fixed. Same typo was also found on P11 L2.

P12, L23: What does the “event-time” correspond to? Is it between NPF start and NPF end, or between NPF start and end of NPF event?

The definition for ‘event-time’ (from NPF start to NPF end) was given in the caption of figure 9, but we now added it to the text as well.

Fig. 1: could the authors change the colour of the red vertical line, which is not easy to distinguish from the background?

We added thin black edges for all of the vertical lines to make them more easily distinguishable.

New particle formation, growth and shrinkage at a rural background site in western Saudi Arabia

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Abstract. Atmospheric aerosols have significant effects on human health and the climate. A large fraction of these aerosols originates from secondary new particle formation (NPF), where atmospheric vapors form small particles that subsequently grow into larger sizes. In this study, we characterize NPF events observed at a rural background site of Hada Al Sham (21.802° N, 39.729° E), located in western Saudi Arabia, during the years 2013–2015. Our analysis shows that NPF events occur very frequently at the site, as 73 % of all the 454 classified days were NPF days. The high NPF frequency is likely explained by the typically prevailing conditions of clear skies and high solar radiation, in combination with sufficient amounts of precursor vapors for particle formation and growth. Several factors suggest that in Hada Al Sham these precursor vapors are related to the transport of anthropogenic emissions from the coastal urban and industrial areas. ~~In Hada Al Sham, the precursor vapors seem to be related to the transport of anthropogenic emissions from the coastal urban and industrial areas, since no NPF events are observed in air masses coming from the sparsely inhabited inland.~~ The median particle formation and growth rates for the NPF days were $8.7 \text{ cm}^{-3} \text{ s}^{-1}$ ($J_{7\text{nm}}$) and 7.4 nm h^{-1} ($GR_{7-12\text{nm}}$), respectively, both showing highest values during late summer. ~~In addition~~ Interestingly, the formation and growth rates increase as a function of the condensation sink, likely reflecting the common anthropogenic sources of NPF precursor vapors and primary particles affecting the condensation sink. ~~large primary particles and NPF precursor vapors.~~ 76 % of the NPF days showed an unusual progression, where the observed diameter of the newly formed particle mode started to decrease after the growth phase. In comparison to most long-term measurements, the NPF events in Hada Al Sham are exceptionally frequent and strong both in

terms of formation and growth rates. In addition, the frequency of the decreasing mode diameter events is higher than anywhere else in the world.

1. Introduction

The effect of atmospheric aerosols on the Earth's radiative balance, via scattering, absorption and cloud interactions, is the single largest factor limiting our understanding of future and past climate changes (Stocker et al., 2013). In addition to the climate effects, aerosols are known to be detrimental to human health, with outdoor particulate pollution being the cause of more than 3 million premature deaths in the year 2010 (Lelieveld et al., 2015). These effects include the contribution of both primary and secondary aerosol particles. Primary particles are emitted into the atmosphere directly as particles, while secondary particles are formed from atmospheric vapors in new particle formation (NPF) events. Measurements of sub-micron particle number size-distributions (PNSDs) have shown that NPF events are a global phenomenon (Kulmala et al., 2004) and they are estimated to produce around half (31–66 %) of the cloud condensation nuclei (CCN) in the lower troposphere (Yu and Luo, 2009; Merikanto et al., 2009; Gordon et al., 2017). Even in polluted regions, where the primary emissions are high, NPF events are found to occur frequently (Yu et al., 2017) and they-is are estimated to be a significant contributor to the particle number concentrations (Yue et al., 2011; Kulmala et al., 2016; Yao et al., 2018). Despite the importance of NPF, many aspects related to the initial formation and subsequent growth of secondary aerosol particles remain unknown. While sulfuric acid is widely regarded as the most important precursor for NPF, it is clear that other compounds are needed to explain particle formation and growth rates in ambient measurements, especially in the boundary layer (Kirkby et al., 2011; Ehrhart et al., 2016). Stabilizing bases, such as ammonia and dimethylamine, low-volatility oxidation products of VOCs (Volatile Organic Compounds) and ions have been shown to enhance particle formation rates and bridge some of the gaps between theoretical evaluations, laboratory studies and ambient measurements (Yu et al., 2012; Almeida et al., 2013; Kürten et al., 2016; Kürten et al., 2014; Kürten et al., 2018; Zhang et al., 2004; Riccobono et al., 2014; Yu et al., 2018). The initial particle-forming compounds will also participate in growing the particles, but in order to explain the observed growth rates, the presence of more abundant condensing (or otherwise particle mass forming e.g. via heterogeneous oligomer formation) species is required (Nieminen et al., 2010; Riccobono et al., 2012). These species are very likely to be low- or semi-volatile organic compounds, formed by the oxidation of either biogenic or anthropogenic VOCs. ~~These species are likely to be photochemically formed low- or semi-volatile oxidation products of either biogenic or anthropogenic VOCs~~ (Smith et al., 2008; Tröstl et al., 2016; Dall'Osto et al., 2018). Overall, the mixture of compounds and the relative importance of different species participating in aerosol formation and growth is expected to vary depending on the ambient conditions; in some coastal environments NPF can be driven by iodine compounds (Sipilä et al., 2016) and e.g. in urban areas the uptake of nitrate can contribute significantly to aerosol mass (Li et al., 2018). Predicting all the occurring interactions in the atmosphere is impossible without observations from several different environments. While PNSD measurements have been conducted in a wide range of environments (Kerminen et al., 2018), especially continuous long-

term measurements are still fairly uncommon and largely focused on Europe and the mid-latitudes (Nieminen et al., 2018). Long-term measurements are needed for obtaining reliable estimates on the average properties and ~~the~~ seasonal tendencies related to the of atmospheric NPF. Such data are essential in evaluating the performance of global models, which currently have large uncertainties in simulating atmospheric NPF, as well as its contribution to CCN budgets and aerosol radiative effects in different environments (Pierce and Adams, 2009; Makkonen et al., 2012; Gordon et al., 2016; Semeniuk and Dastoor, 2018). ~~which are important for e.g. model validation.~~

Recently, several NPF studies have pointed out an interesting phenomenon, where the average diameter of the particle mode formed in an NPF event begins to decrease after the growth phase. This is often referred to as aerosol shrinkage, but we will use the term DMD (decreasing mode diameter) event, since aerosol shrinkage quite directly implies a reduction in the size of individual particles, which is not necessarily the case. Such DMD events have been observed especially in subtropical regions (Yao et al., 2010; Backman et al., 2012; Cusack et al., 2013; Young et al., 2013; Zhang et al., 2016; Alonso-Blanco et al., 2017), but also in the temperate climate (Skrabalova, 2015; Salma et al., 2016). Typically, these DMD events are suggested to be caused by the evaporation of semi-volatile compounds, due to changes in environmental conditions. However, the reduction in the mean diameter of a particle mode may also occur without evaporation, if smaller particles are transported to the measurement site (Kivekäs et al., 2016).

In this paper, we study the aerosol particle number-size distribution dynamics at Hada Al Sham, Saudi Arabia, during February 2013–~~March~~–February 2015. The environment is quite unique due to high anthropogenic and low biogenic emissions, and a distinct segregation between the surroundings in different directions. To our knowledge, these are the first comprehensive long-term aerosol measurements conducted in the Arabian Peninsula. Two articles from the same measurement campaign are already published, describing the aerosol physical (Lihavainen et al., 2016) and optical properties (Lihavainen et al., 2017). The former article also contains analysis of particle number concentrations, which will not be presented in this study. This work focuses on identifying and characterizing NPF in the study region. We show that the NPF events in Hada al Sham are, in comparison to the locations analysed in the exhaustive study by Nieminen et al. (2018), exceptionally frequent and intense, in terms of both particle formation and growth rates. We also make a detailed investigation of the typical diurnal cycle related to the NPF events and determine how the events are impacted by different environmental variables, including meteorological conditions and the concentrations and sources of primary aerosol particles and aerosol precursor compounds.

2. Measurements and methods

2.1 Measurement site and instrumentation

Hada Al Sham (21.802 °N, 39.729 °E, 254 m a.s.l.) is a small city in western Saudi Arabia (see Fig. 3). There are no major sources of anthropogenic emissions in the immediate vicinity and the site can be described as a rural background site.

5 Biogenic emissions are also presumably minor due to the arid desert climate and lack of vegetation. Sparsely inhabited desert-like areas cover the inland in the N-SE direction from the measurement site, but the coastal regions in the western sector are densely populated. Jeddah, the second largest city in Saudi Arabia, is located by the Red Sea ~ 60 km to the west from Hada Al Sham. In ~~this region~~ Jeddah, there are several major emission sources, including power generation plants, a sea water desalination plant, as well as an airport and a harbor, both of which experience heavy traffic due to the combined effect
10 of economic growth and the vicinity of Makkah located ~ 40 km to the south from Hada Al Sham. Other major emission sources in the region include a petrochemical refinery and steel industry in Rabigh (~ 130 km NW, see Fig. 3), and a large oil refinery in Yanbu (~ 290 km NW). e.g. an oil refinery, a sea water desalination plant and a power generation plant. There is also an airport and a harbor that both experience heavy traffic due to the economic growth and the vicinity of Makkah that is located ~ 40 km to the south from Hada Al Sham. The densely populated coastal region and the sparsely populated inland
15 are separated by mountains running along the coast of the Red Sea. Hada Al Sham is located by the western slopes of these mountains and is thus topographically more connected to the coastal region.

The measurements were conducted at the Agricultural Research Station of King Abdulaziz University from ~~November~~ February 2012–2013 to February 2015. The instruments were placed inside a container, located on a sand field, with a
20 distance of ~ 100 m to the nearest trees and other obstacles. The temperature was kept stable inside the container at ~ 25 °C. The sample air inlets were located at a height of 4–4.5 m and the sample air flow rate was 16.7 l min⁻¹. A more thorough description of the measurement setup and the used instruments can be found in Lihavainen et al. (2016). In this study, we focus on the PNSD measurements in the mobility diameter range of 7–850 nm, obtained using a twin DMPS (Differential Mobility Particle Sizer (Wiedensohler et al., 2012)), and the meteorological parameters (temperature, relative humidity, wind
25 speed and wind direction), which were measured with a Vaisala WXT weather station. The twin DMPS used here consists of a short and a medium Hauke-type DMA (Differential Mobility Analyzer, custom-made) and two CPCs (Condensation Particle Counter, TSI 3772). To study the effects of mineral dust, we utilized the PNSDs in the aerodynamic diameter range of 0.5–10 µm, measured with an APS (Aerodynamic Particle Sizer, TSI 3321), and the mass of particles smaller than 10 µm (PM₁₀), measured with a beta hybrid mass monitor (Thermo Scientific 5030). The inlet leading to the DMPS and the APS
30 was equipped with a PM₁₀ filter and a twin diffusion dryer, which kept the sample air relative humidity mainly below 50 %. PNSD measurements were corrected for diffusional losses in the inlet line and they were at maximum around 15 % at the lower size limit of the DMPS. The PM₁₀ measurements were made from a separate inlet, equipped with a standard heater for sample drying. Since no gas phase measurements were conducted during this campaign, we used data from the Ozone

Monitoring Instrument (OMI) on-board NASA's Aura satellite (Levelt et al., 2006) to estimate the SO₂ concentrations (OMI Level 2 SO₂ Planetary Boundary Layer product (Li et al., 2013)) at Hada Al Sham and its surroundings.

2.2 NPF event classification

NPF event classification was done for the measurement days based on the visual interpretation of PNSDs, as described by Dal Maso et al. (2005). Each day was classified as either: (1) NPF day, (2) non-event day or (3) undefined. In short, a day is classified as an NPF day if a new growing mode of particles appears in the nucleation mode ($d_p < 25$ nm) and the growing mode is observed to persist for several hours (see Fig. 1 for an example of an NPF day and Fig. A1 for time series of PNSDs from selected periods). If some, but not all, of the above-mentioned criteria are fulfilled, the day is classified as undefined. In addition, a day is classified as undefined if the time development of the newly formed particle mode is highly erratic, or if the mode is not continuous due to significant breaks. Only the days, when there is clearly no indication of NPF, are classified as non-event days.

As an addition to the traditional classification, each NPF day was further classified based on whether the mean diameter of the mode formed in the NPF event clearly starts to decrease after the growth phase (see Fig. 1), or not. These days are referred to as DMD (Decreasing Mode Diameter) days and non-DMD days, respectively.

2.2.1 Event times

To describe the progression of the NPF events, we determined the points in time when: (1) NPF is first observed in the smallest size-bins of the DMPS measurements, (2) NPF is no longer observed, (3) the mode diameter of the newly formed particles starts to decrease and (4) the mode formed in the NPF event is no longer distinguishable from the background aerosols due to decreased number concentrations or e.g. changes in air masses. These times were determined visually from the PNSDs and they are referred to as NPF start, NPF end, DMD start and NPF event end, respectively (Fig. 1). NPF start times were only determined for the days when ~~the~~-NPF was observed all the way from the smallest size bins, and the NPF event end times only for the days when the event ended during the same day as ~~the~~-NPF had started.

2.3 Determination of particle formation and growth rates

The formation and growth rates of newly formed particles are important quantities in describing NPF (Kulmala et al., 2012). They provide information on the strength of the NPF events and are closely connected to the atmospheric factors driving the process, such as the concentrations of condensable vapors. The growth rates of the freshly formed particles have a critical role in their probability to survive into climate-relevant sizes, as particles that grow too slowly are removed by coagulation with larger pre-existing particles (Kuang et al., 2009).

30

In this study, the particle growth rates were determined by following the time development of the geometric mean diameters (GMDs) obtained from log-normal fits to the PNSD at each measurement time. The fitting was done using an automatic algorithm developed by Hussein et al. (2005), which analyses the measured PNSD, fits 2–3 log-normal modes and returns the fitting parameters. In practise, the growth rates are determined by plotting the ~~fitted mode diameters~~ GMDs of fitted modes together with the PNSD and making a linear fit to the points selected to represent the mode formed in the NPF event (Fig. 1). Now, the diameter growth rate (GR) in the size range Δd_m can be calculated simply as the slope of the fitted line:

$$GR_{\Delta d_m} = \frac{\Delta d_m}{\Delta t} \quad (1)$$

where Δd_m is the change in the geometric mean diameter during the time interval Δt . The growth rates presented in this study were determined so that they would best describe the growth of particles in the diameter range 7–12 nm, as this is the range used in the determination of the formation rates. In principle, this could be done by ~~always only~~ selecting ~~only~~ the ~~mode fit~~ GMD points below 12 nm, but since single points are subject to fluctuations in the PNSD, and the number of fitted points below the 12 nm threshold was often quite small, points above 12 nm were ~~also~~ often included into the fit to obtain a more robust estimate of the particle growth. This was, however, done only if the growth rate seemed to remain constant above the 12 nm threshold.

The formation rate (J) of particles with diameter d_p can be determined using the equation (Kulmala et al., 2004):

$$J_{d_p} = \frac{dN_{\Delta d_p}}{dt} + CoagS_{\Delta d_p} N_{\Delta d_p} + \frac{GR_{\Delta d_p}}{\Delta d_p} N_{\Delta d_p} \quad (2)$$

where $N_{\Delta d_p}$ is the number concentration and $CoagS_{\Delta d_p}$ is the coagulation sink (Sect. 2.4) of the particles in the size range Δd_p . Therefore, the formation rate is defined as the flux of particles past the lower limit of the size range, and it is obtained by adding up the change in the observed particle number concentration with the losses of particles due to coagulation and growth out of the size range. In this study, we calculated the formation rate of 7 nm particles using the diameter range 7–12 nm.

2.4 Calculation of coagulation and condensation sinks

The coagulation sink describes the rate, at which particles are lost due to collision and coalescence with larger particles. The collisions can occur due to differing settling velocities, turbulence, electric interactions or Brownian motion. However, when describing the coagulation of submicron particles in typical atmospheric conditions, the coagulation due to Brownian motion is by far the most significant. When only this mechanism is considered, the coagulation sink can be calculated by integrating over the PNSD (Kulmala et al., 2001):

$$CoagS_{d_p} = \int K(d_p, d'_p) n(d'_p) dd'_p = \sum_{d_{p,i}=d_p}^{d_{p,max}} K(d_p, d'_{p,i}) N_{d'_{p,i}} \quad (3)$$

where $K(d_p, d'_p)$ is the coagulation coefficient (Seinfeld and Pandis, 1998; Fuchs, 1964), which describes the probability of particles with diameters d_p and d'_p to collide. This coefficient is proportional to the particle surface area and it increases with increasing size difference between the colliding particles. Both the coagulation sink of nucleation mode particles and the condensation sink are largely determined by the Aitken and accumulation mode particles that typically dominate the total particle surface area. The coagulation and condensation sinks were calculated using the DMPS measurements, which sets the upper diameter limit ($d_{p,max}$) of particles included in the calculation at 850 nm. In the calculation of the coagulation sink, we used the geometric mean of the diameter range (7–12 nm) to approximate the size of coagulating particles.

The condensation sink (*CS*) describes the ability of the aerosol population to remove condensable vapors from the atmosphere. The concept is analogous to the coagulation sink, defined with Eq. 3, but now instead of the particle loss rate, the rate at which vapors condense on to pre-existing aerosol particles is considered. Similarly to the coagulation sink, the *CS* is calculated by integrating over the PNSD (Kulmala et al., 2001):

$$CS = 2\pi D \int d_p \beta_m(d_p) n(d_p) dd_p = 2\pi D \sum_{d_{p,i}=d_{p,min}}^{d_{p,max}} \beta_m(d_{p,i}) d_{p,i} N_i \quad (4)$$

where D is the diffusion coefficient of the condensing vapor and β_m is the correction coefficient for the transitional regime (Fuchs and Sutugin, 1970). Here, the *CS* is calculated using the properties of sulfuric acid as the condensing vapor. In practice, this means that the *CS* describes the loss rate of such vapors that condense irreversibly onto the particles upon each collision.

2.5 Air mass history

Air mass history was studied by calculating particle retroplumes using a Lagrangian particle dispersion model FLEXPART (FLEXible PARTicle dispersion model) version 9.02 (Stohl et al., 2005). ECMWF (European Centre for Medium-Range Weather Forecast) operational forecast with 0.15° horizontal and 1 h temporal resolution was used as the meteorological input into the model. The number of height levels in the meteorological data was 91 before 25 June 2013 and 137 after that.

The model was run for a time period from February 2013 to May 2014. During this period, a new release of 50 000 model particles, distributed evenly between 0–100 m above the measurement site, occurred every 1 hour. The released particles were traced backwards in time for 72 h, unless they exceeded the model grid (0–45°N, 15–70°E, resolution: 0.05°). The model time step was 10 min, but in the calculation of turbulent wind components, the use of a shorter time step, determined internally by the model, was allowed. The parametrization of moist convection was also set on to improve the quality of the model run. Particle wet and dry deposition were not considered, since in this work, the model is only used to study the movements of the air masses.

The model output was saved every 1 hour and, in relevance to this study, it contains: (1) the emission sensitivity field i.e., a matrix whose values are proportional to the time the model particles have spent over each grid point during the last hour and (2) a point of the average trajectory that is determined by cluster analysis (Seibert and Frank, 2004) from the locations of the model particles. In addition, we use the atmospheric boundary layer (ABL) height at Hada Al Sham, obtained from the
5 ECMWF operational forecast.

3. Results and discussion

3.1 NPF event frequency

The results of the NPF event classification are presented in Fig 2. The total NPF event frequency was found to be very high, as 73 % of all the classified days (454) were NPF days. Out of the NPF days, 76 % were DMD days, meaning that only
10 about one quarter of the NPF events showed monotonic growth, which is the typical progression of NPF in most environments. Only 4 % of the days were classified as non-events leaving 23 % as undefined. It should be noted, that the majority of the undefined days were likely days when new particles were formed, but continuous growth of particles was not observed at Hada Al Sham due to unfavorable meteorological conditions (referred to as 'failed events' by -Buenrostro Mazon et al. (2009)).~~showed some features representative of NPF events but these features were not clear enough for the~~
15 ~~days to be classified as NPF days.~~

The monthly fractions in Fig. 2 show that the NPF frequency is high (55–85 %) throughout the year and that no clear seasonal pattern is observed. This implies that the NPF events ~~at this site~~ are not limited by any factor with a strong seasonal variability. The most notable deviations from the average frequency are found in June, November and December, which all
20 have a higher than average fraction of non-event days. Although no seasonal cycle is seen in the total NPF frequency, the DMD events are more frequent during the summer (and autumn) months and less frequent during winter. The fraction of DMD events from all NPF days is highly variable, ranging from 33 % in November to 95 % in September.

The average NPF event frequency of 73 % in Hada Al Sham is among the highest event frequencies obtained from ~~long~~
25 ~~term~~long-term measurements. Nieminen et al. (2018) compared PNSD measurements, consisting of at least one full year of data, conducted at 36 different sites around the world. They observed that NPF events are most frequent in South Africa, where the NPF frequencies from three different sites were 69, 75 and 86 % (Hirsikko et al., 2012;Vakkari et al., 2011;Vakkari et al., 2015). Thus, the NPF fraction of 73 % obtained from the measurements presented here, would take the third place in this global comparison. The high NPF event frequency is a direct indication of typically favorable conditions
30 for new particle formation and growth. NPF event frequency has been shown to be affected by at least: solar radiation, SO₂ concentration, vapor and particle sinks and air mass origins (Nieminen et al., 2015;Kerminen et al., 2018). The effects of these factors are discussed in the following paragraphs in order to explain the observed NPF frequency in Hada Al Sham.

The connection between solar radiation and NPF events is related to atmospheric photochemistry: the production of sulfuric acid, which is ~~widely~~ regarded as the driving compound of atmospheric new particle formation in most environments (Weber et al., 1997; Birmili et al., 2003; Kuang et al., 2008; Paasonen et al., 2010; Yao et al., 2018), occurs mainly via oxidation of SO₂ by OH, and the concentration of OH is dependent on the intensity of ~~the~~ solar radiation. Baranzadeh et al. (2014) and Dada et al. (2017) studied the effect of cloudiness on the NPF frequency at ~~the~~ SMEAR II station in Hyytiälä, Finland, and showed that the NPF frequency increased from ~ 0 % to over 50 % with decreasing cloudiness. During some months, the NPF frequency on clear sky days reached over 70 % also in Hyytiälä, which clearly demonstrates the importance of photochemistry. The exact number of clear sky days in Hada Al Sham could not be determined as no radiation measurements were made, but in general the radiation conditions in this area (Alnaser et al., 2004) are highly favorable for the occurrence of NPF.

Solar radiation alone is, of course, not sufficient to cause NPF if no precursor vapors for the production of nucleating and condensing compounds are available. ~~The~~ SO₂, required for the production of sulphuric acid, is emitted especially from traffic and industries that process or consume fossil fuels. In Hada Al Sham, there are no significant sources of SO₂ but there are plenty of sources within a 100 km radius of the site, which includes the large urban and industrial areas of Jeddah and Makkah. Figure 3 displays the average SO₂ concentration, retrieved by satellite measurements, in the surroundings of Hada Al Sham on NPF days. Even though these concentrations represent the amount of SO₂ in the whole vertical column, the values should reflect the concentrations in the boundary layer since the most significant SO₂ sources are located at ~~the~~ ground level (for reference, 1 DU corresponds to 11 ppbv when distributed into a 1 km boundary layer at T = 300 K and P = 1 atm). The figure shows that the SO₂ concentrations on the coastal region are very high, close to those obtained from the most SO₂-polluted regions in the world (Krotkov et al., 2016). The horizontal transport of the emissions seems to be restricted by the mountains, which causes the accumulation of the concentrations on a relatively narrow area and creates a distinct boundary in the SO₂-levels compared to the inland. The observed high SO₂ concentrations, in combination with the radiation conditions, suggest that the production rate of sulfuric acid is ~~also~~ high at this site.

In Fig. 4, we compare the air mass history during the mornings of NPF event and non-event days. The shown emission sensitivities are calculated from the 24-hour retroplumes, initiated at the time when NPF is typically taking place (10:00 LT). The comparison shows significant differences in the air mass origins between these cases. On NPF days (Fig. 4a), the air masses observed in Hada Al Sham originate mainly from a narrow strip that extends along the coast and includes the regions of significant SO₂ emissions (see Fig 3). The formation of such air mass source region can be explained by the typically prevailing large-scale winds that blow along the Red Sea, due to tunnelling caused by the steep shores. Over the coastal cities, anthropogenic emissions are introduced into the air masses and then transported to Hada Al Sham with the sea breeze that usually starts to develop already early in the morning. On non-event days (Fig. 4b) there are no signs of the sea breeze or

the typically prevailing large scale wind. Instead, the source regions initially point towards east, after which they spread to cover larger areas in the inland. The inland is sparsely inhabited and the sources of anthropogenic emissions are few, which can also be seen as lower average SO₂ concentrations in the surroundings of Hada Al Sham on non-event days than on NPF days (Fig. A2). The fact that all of the non-event days are observed when the air masses are coming from the inland ~~implies~~ suggests that the NPF events observed in Hada Al Sham ~~might depend on~~ be related to the emissions from the coastal anthropogenic activities. In addition, the influence of marine air could be beneficial for NPF occurrence e.g. due to lower condensation sink or some specific precursors.

The markedly different wind conditions between the NPF and non-event days can also be seen from the local measurements at Hada Al Sham (Fig. 5a and b). On NPF days, the weak nocturnal easterly wind (land breeze) turns westerly and its speed starts to increase between 8:00 and 10:00 due to the development of the sea breeze. This shifts the air mass source regions to the coastal areas, as seen in Fig. 4a. On non-event days, such change in the wind direction is not seen. During the night and early morning, the wind is easterly, similar to the NPF days, but on non-event days, the easterly wind is significantly stronger. This inhibits the development of the sea breeze circulation and the westerly wind associated with it.

The strong easterly winds on non-event days seem to resuspend dust from the inland desert, which can be seen as the simultaneous increase of PM₁₀ with the wind speed (Fig. 5d). Both the wind speed and the PM₁₀ obtain their largest values around the same time when NPF typically starts (see Fig. 6). It is therefore possible that, in addition to the fewer emission sources in the inland, the NPF is inhibited by the wind-blown dust, which can reduce the concentrations of the clustering and condensing vapors by limiting ~~the~~ solar radiation, and by acting as a sink for reactive gases and oxidants (Hanisch and Crowley, 2003; Usher et al., 2003). The condensation sink does not, however, seem to be an determining-inhibiting factor in thefor NPF occurrence, as significantly larger CS values are observed on NPF days than on non-event days (Fig. 5c). Here the presented CS values are calculated including only the particles measured by the DMPS ($d_p < 850$ nm), but including the larger particles measured by the APS (d_p up to 10 μ m) did not have a significant effect on these results (~~not show~~ hereincrease in CS on non-event days was around 10 %). Furthermore, recent studies have shown indications that mineral dust, mixed with anthropogenic emissions, could actually enhance NPF due to heterogeneous production of sulfates and hydroxyl radicals (Nie et al., 2014; Xie et al., 2015). In Hada Al Sham, the concurrence of high PM₁₀ values and non-event days would then mainly highlight the lack of anthropogenic emissions in the inland, although we cannot quantify the effect of PM₁₀ on radiation.

Higher CS values during NPF days (Fig. 5c) are also reported from a high altitude site in the Swiss Alps (Boulon et al., 2010). Here this connection is speculated to stem from the coupled appearance of NPF precursor vapors and CS due to their common lower altitude sources. Similar situation could apply to Hada Al Sham in case both the CS and the NPF precursor

vapors originate mainly from the same sources. The high CS during the calm nights and early mornings of NPF days could then suggest enhanced accumulation of precursor vapors, thus facilitating the occurrence of the NPF process after sunrise.

3.2 NPF event progression and characteristics

Figure 6 displays the frequency histograms of the points of time describing the progression of the NPF events (see Sect. 2.2.1), together with the diurnal variation of meteorological parameters and CS. In Fig. 7, the NPF progression times are plotted for a year-long period from June 2013 to June 2014 to show their seasonal variation. NPF events are typically observed to start slightly before 9:00 (Fig. 6a). On a seasonal scale, the starting times change according to the changes in the time of sunrise (Fig. 7). This observation highlights the importance of photochemistry for the NPF, especially, since none of the NPF events start before sunrise. There is, however, quite significant day-to-day variation in the starting times, which is not explained by differences in the times of sunrise. Some of this variation can be attributed to differing growth rates between the NPF events. This is because the NPF starting times are defined here as the times when new particles are observed at the size of 7 nm, even though the formation of new particles actually starts from the molecular scale. Therefore, the time that it takes for the small particles/molecular clusters of the size $\sim 1\text{--}2$ nm (Kulmala et al., 2013) to grow and reach the lower limit of the DMPS, affects the starting times shown here. In some cases, the later NPF starting times are caused by a delayed shift in the wind direction from the inland to the coastal side of the measurement site (Fig. A3). This is in agreement with the interpretation that the NPF events observed in Hada Al Sham ~~depend on~~ are related to the transport of emissions from the coastal regions. The connection between the shift in the wind direction and the onset of NPF can also be seen from the average values, as the typical onset time of the NPF events coincides with the onset of the westerly sea breeze (Fig. 6a and c). In addition to the radiation conditions and the wind direction, the starting time seems to be connected to a drop in the CS, which is likely caused by the increasing ABL height (Fig 6a and b).

The formation of new particles lasts, on average, for about 3 hours ending around noon. The NPF end times are possibly affected by the simultaneously increasing CS (Fig. 6a and b), which is caused by the growth of both the freshly formed and the pre-existing accumulation mode particles into larger sizes. With increasing CS, ~~the~~ NPF precursor vapors and ~~the~~ new small particles are more likely to end up contributing to the growth of the pre-existing aerosol, rather than forming a new growing mode of their own. The end times could also be linked to a weakening production of condensable vapors, as discussed in the following paragraph in the context of the DMD events.

On the majority of the NPF days, the particle growth phase is followed by a DMD event. Typically, the DMD phase starts in the afternoon around 15:00, approximately 6 hours after ~~the~~ NPF start (Fig. 7). The onsets of the DMD events are seemingly concurrent with the maxima of the wind speed, ABL height and the temperature. This could indicate that the DMD events are caused by particle evaporation, which is triggered by the increased saturation vapor pressure at elevated temperatures and the dilution of vapor concentrations due to the ABL development and wind-induced mixing. This could also explain the

observed summer maximum in the DMD event frequency, since these variables, which are likely to promote particle evaporation, obtain their largest values during summer. In addition, ~~the~~ particle evaporation might be facilitated by the decreased photochemical production of condensable vapors after the maximum intensity of solar radiation.

5 A vast majority (~ 85 %) of the NPF events observed in Hada Al Sham end during the same day they started, approximately 3 hours after ~~the~~ DMD start or 9 hours after ~~the~~ NPF start (Fig. 7). Here the NPF event end times were defined as the points in time, when the number concentrations, associated to the particle mode formed in the NPF event, drop significantly, making the mode indistinguishable from the background aerosols (see Fig. 1). Based on the observations presented earlier in this work, explaining the ending of the NPF events would be rather straightforward if they were related to the wind turning easterly at around 23:00 (Fig. 6c). However, the NPF events are typically observed to end hours before this (Fig. 6a), while the winds are still westerly. This indicates, that also in the westerly direction, an area where NPF is not occurring is reached. When the air masses that resided in this area during the active NPF hours (NPF start–NPF end) are transported to the measurement site, the particle mode related to NPF is no longer seen. An order-of-magnitude estimate for the westerly extent of the NPF area can be obtained by multiplying the average duration of the NPF events by the average wind speed: $10 \text{ h} * 10 \text{ km h}^{-1} = 100 \text{ km}$. This is comparable with the distance from Hada Al Sham to the coast of the Red Sea where the concentration of SO_2 drops rapidly (Fig. 3). Therefore, it seems that no NPF is happening outside the region of strong contribution from anthropogenic emissions. It is reasonable to assume that there is no discrete boundary between the regions where NPF is and is not occurring. Instead, when moving away from the region of high emissions, particle formation and growth rates can be expected to decrease gradually due to the decrease in the concentrations of participating vapors. This provides another possible explanation, in addition to particle evaporation, for the observed DMD events: when the particles formed in less favourable conditions are transported to the measurement site, it is possible that they would have grown less than the previously observed particles. Continuous observations of particles that have grown less and less can produce a DMD event without particle evaporation or other shrinkage (Kivekäs et al., 2016). The cause of the DMD events observed in Hada Al Sham is further investigated in a future study.

25 3.3 Particle formation and growth rates and their dependence on environmental conditions

Figure 8 shows the seasonal variation of the particle formation rates ($J_{7\text{nm}}$) and growth rates ($GR_{7-12\text{nm}}$) determined for the NPF events observed in Hada Al Sham. The formation rates (Fig. 8a) vary mostly between 1 and $50 \text{ cm}^{-3} \text{ s}^{-1}$, having an annual median (determined as the median of all observations) of $8.7 \text{ cm}^{-3} \text{ s}^{-1}$. The variation in the growth rates (Fig. 8b) is slightly smaller with the values ranging mainly from 2 to 20 nm h^{-1} and having a median of 7.4 nm h^{-1} . The variation of these values is similar to the observations from several locations around the world (Kulmala et al., 2004) but the median values are very much in the high end of observations (Nieminen et al., 2018; Kerminen et al., 2018). Nieminen et al. (2018) reported, in their global study, the highest annual average particle formation rate from Beijing ($J_{10\text{nm}} \sim 7 \text{ cm}^{-3} \text{ s}^{-1}$) and found an overall increasing trend of formation rates with increasing degree of anthropogenic influence.

Both the formation and growth rates show a similar seasonal cycle, with largest values during the summer and early autumn (Fig. 8a, b). For the growth rates, a summer maximum is often observed also globally, while the formation rates peak in most locations during spring (Nieminen et al., 2018). The different seasonal cycles of J and GR at these sites could result from different species controlling the initial formation and further growth of the particles. In many places, the summer maximum in growth rates might be related to increased photochemical activity and increased emissions of biogenic VOCs as a function of temperature (Yli-Juuti et al., 2011), while the formation of particles could be more dependent on anthropogenic emissions (Nieminen et al., 2014). Conversely, the similar cycles of both J and GR in Hada Al Sham could then suggest that here the precursor sources of particle-forming and -growing vapors are similar. Since the emissions from biogenic sources are expected to be minor, due to the lack of vegetation in the area, these sources are likely to be anthropogenic. In Fig. 8a, we show the monthly crude oil consumption in Saudi Arabia which seems to peak around the same time as J and GR. Thus the increased emissions from energy production, resulting from the growing need for air conditioning during summer, could possibly explain the observed seasonality in J and GR, although increased solar radiation is also likely to play a role. Both the formation and growth rates show a similar seasonal cycle, with largest values during the summer and early autumn (Fig. 8a, b). For the growth rates, a summer maximum is often observed also globally (Nieminen et al., 2018), which is possibly explained by increased photochemical activity and increased emissions of biogenic VOCs (BVOCs) as a function of temperature (Yli-Juuti et al., 2011). In Hada Al Sham, the increased photochemistry is also likely to play a role, but we expect the influence of BVOCs to be minor, simply due to the lack of vegetation in this area. Here, the seasonal cycle of the formation and growth rates is, however, likely affected by the increased emissions from energy production during summer (see crude oil consumption in Fig. 8a), which results from the growing need for air conditioning.

In Fig. 8, the formation and growth rates are presented separately for the DMD and non-DMD NPF events. The comparison between these two cases is difficult due to the small number and uneven distribution of the non-DMD events. Regardless, during November–January, when the number of non-DMD events is especially large, the growth rates on non-DMD days are quite consistently lower than those on the DMD days (Fig. 8b). This would imply that the conditions between these cases are different already in the early stages of the NPF events, even though the DMD phase does not occur until hours later. The difference in the formation rates (Fig. 8a) is, however, less pronounced. One possible explanation for this could be higher concentrations of some semi/intermediate-volatility compounds on DMD days that would not participate in the initial particle formation, but would gain effectiveness with increasing particle size, due to decreasing Kelvin effect.

In Fig. 9, the formation rates are presented as a function of the relevant meteorological variables (a–d), PM_{10} (e) and CS (f) (see Fig. A4 for growth rates). Despite the rather pronounced seasonal variation in the formation rates (Fig. 8a), no correlation is found between the event-time (from NPF start to NPF end) temperature and J, when all data is considered (Fig. 9a). This is likely due to the temperature peaking already in July, while the formation rate peaks later in August.

Interestingly, a strong negative correlation between J and T ($r = -0.62$, $p = 3 \times 10^{-5}$) is observed when considering only the summer months (MJJA). This might reflect the negative effect of increasing vapor volatility ~~for~~on particle formation.

Figure 9b shows that the formation rate increases with increasing RH. This is the expected relationship between these two variables (Almeida et al., 2013; Duplissy et al., 2016; Kürten et al., 2016), as water vapor is known to participate in the cluster formation with sulfuric acid. However, in ambient measurements such correlation could be caused by processes that are not necessarily related to the RH effect itself. For example, here the higher RH could be related to the coastal origin of the air masses, and simultaneously to higher anthropogenic emissions from the coastal sector. To examine this, the correlation was calculated separately for winds coming from the S–W sector, where the SO₂ distribution seems most uniform (Fig. 3), but similar relationship to the case with all data was found. Furthermore, the RH relationship does not seem to be related to the seasonal variation, as both high and low RH values are observed throughout the year.

Higher formation rates seem to be favored by low wind speed and low ABL height (Fig. 9c and d). During low ABL conditions, the near-surface anthropogenic emissions are distributed into a smaller volume, which could then lead to higher vapor concentrations and particle formation rates. Analogously, the accumulation of emissions, per unit volume of air, from a spatially limited emission source area is increased during low wind speed conditions. Interestingly, the lowest event-time ABL heights are observed during the summer months (Fig. 9d), meaning that in addition to the summer NPF events occurring earlier in the absolute sense (see Fig. 7), they also take place earlier with respect to the boundary layer development. This observation could be caused by the higher emissions during summer (see Fig. 8a), since if the concentrations of NPF precursors are higher, the onset of NPF events is likely to be more sensitive to an increase in solar radiation. We note that the RH dependence might also arise partly from higher RH in low ABL and wind speed conditions. However, since the correlation of RH and either ABL height or wind speed is weaker than that between RH and the formation rate, we expect this connection to be of secondary importance.

In the end of Sect. 3.1, we discussed briefly the possible interactions between NPF and mineral dust, which is likely the major component of PM₁₀ in Hada Al Sham. We stated that mineral dust can possibly either weaken (increasing sink and decreased solar radiation) or enhance (increasing production of hydroxyl radicals and sulfates via heterogeneous reactions) the NPF events. The enhancing effect has been observed specifically in situations where ~~the~~ mineral dust is mixed with anthropogenic pollution (Nie et al., 2014), which could correspond to the situation in Hada Al Sham. Despite this, no clear correlation was found between PM₁₀ and the formation (or growth) rates (Figs. 9e and A4e). It should be noted that in the case presented by Nie et al. (2014), the time scale of the process is significantly longer, as both the sources of the dust and the anthropogenic emissions are located further away from the measurement site. This allows for a longer interaction time between the dust and the emissions, which might be crucial for the enhancing effect to occur.

Out of the included variables, the strongest correlation is found between particle formation rates and CS (Fig. 9f). This positive correlation is quite interesting, since the concentrations of vapors participating in ~~the~~ NPF are expected to decrease with increasing CS due to their faster loss rate. However, this is generally valid only if the sources of the CS and the condensing vapors are independent from one another. Here, this is likely not the case, but instead the increasing CS presumably represents increasing contribution from the (primary) anthropogenic emissions and is therefore simultaneously linked to higher concentrations of NPF precursor vapors. This is supported by the observation that both the CS and SO₂ concentration are higher during the NPF days than non-NPF days (Figs. 5c and A2).

Figure A4 shows the particle growth rates as a function of the same variables as the particle formation rates in Fig. 9. Overall, the correlations are qualitatively similar, but weaker in the case of GRs. Similarly to the formation rates, the strongest correlation is found with the CS ($r = 0.37$, $p = 2.2 \times 10^{-5}$). The most notable difference is that in the case of GRs, a weak positive correlation with temperature ($r = 0.19$, $p = 0.025$) is observed.

4. Summary and conclusions

The analysis of the aerosol number-size distribution measurements showed that NPF events are a highly frequent phenomenon in Hada Al Sham, with the fraction of NPF days accounting for 73 % of all the classified days. The high NPF frequency is likely explained by the high production of NPF precursor vapors, especially sulfuric acid, in the transported emission plumes from the coastal cities and industrial areas during the typically prevailing ~~cloud~~-~~cloud~~-free and high solar intensity conditions. The fraction of non-NPF days was only 6 % and these days were shown to be linked to strong easterly winds that block the development of the sea breeze, which typically brings the polluted air masses to Hada Al Sham.

Most of the NPF events in Hada Al Sham displayed an unusual progression, where the diameter of the particle mode related to the NPF event started to decrease after the growth phase. Similar DMD events have been observed in other measurement sites as well, but in Hada Al Sham the frequency of these events was found to be exceptionally high (76 % of all NPF days). The DMD events were more frequent during the summer, and the average onset time of the DMD events was during the afternoon, approximately 6 hours after NPF start.

The median particle formation and growth rates associated with the NPF events were $8.7 \text{ cm}^{-3} \text{ s}^{-1}$ ($J_{7\text{nm}}$) and 7.4 nm h^{-1} ($GR_{7-12\text{nm}}$), respectively. These values correspond to those typically obtained from polluted urban measurement sites. Both the formation and growth rates showed the highest values during summer and autumn months, presumably due to the increased emissions from energy production and the effect of stronger solar radiation on the rate of photochemical reactions. The formation rates were found to obtain higher values in calm conditions where both the wind speed and the ABL height were low and the relative humidity was high. Under such circumstances, anthropogenic emissions are likely to spread and

accumulate throughout the coastal zone, including Hada Al Sham. Both the formation and growth rates obtained higher values in conditions of high CS, which is likely associated with the common anthropogenic sources of NPF precursor vapors and the large primary particles that control the CS.

5 Overall, the findings of this study highlight the importance of anthropogenic emissions and photochemistry for NPF. Due to the transportation of emissions from urban and industrial areas, NPF events were found to be very frequent in Hada Al Sham, located tens of kilometers away from the major sources. The frequency and strength of NPF observed here implies that NPF events might contribute significantly to the budget of both ultrafine and CCN particles, making their health and climate effects relevant topics for further studies in this region. The local conditions at Hada Al Sham, with high levels of
10 regional anthropogenic emissions but presumably low concentrations of biogenic vapors, also allows us to research anthropogenic NPF events in detail. However, further experiments with broader spectrum of instruments are required for determining the vapors responsible for new particle formation and growth, as well as the underlying reasons for the occurrence of the DMD events.

Data availability

15 Data used in this study are available from the corresponding author upon request (simo.hakala@helsinki.fi).

Author contribution

Mamdouh Khoder, Hisham Al-Jeelani, Mansour A. Alghamdi, Heikki Lihavainen, Antti-Pekka Hyvärinen and Tareq Hussein coordinated the measurement programme and carried it out with Kimmo Neitola, [Lubna Dada](#), Ahmad S. Abdelmaksoud, Ibrahim I. Shabbaj and Fahd M. Almeahmadi. Ahmad S. Abdelmaksoud, Ibrahim I. Shabbaj and Fahd M.
20 Almeahmadi also provided resources and data curation. Ville Vakkari provided the essential means for calculating and analyzing the air mass trajectories. Anu-Maija Sundström produced the figures related to the OMI SO₂ data. Simo Hakala and Pauli Paasonen performed the data analysis while Ville Vakkari, Veli-Matti Kerminen, Jenni Kontkanen and Markku Kulmala contributed to the interpretation of the results. Simo Hakala wrote the manuscript with contributions from all co-authors.

25 Competing interests

The authors declare that they have no conflict of interest.

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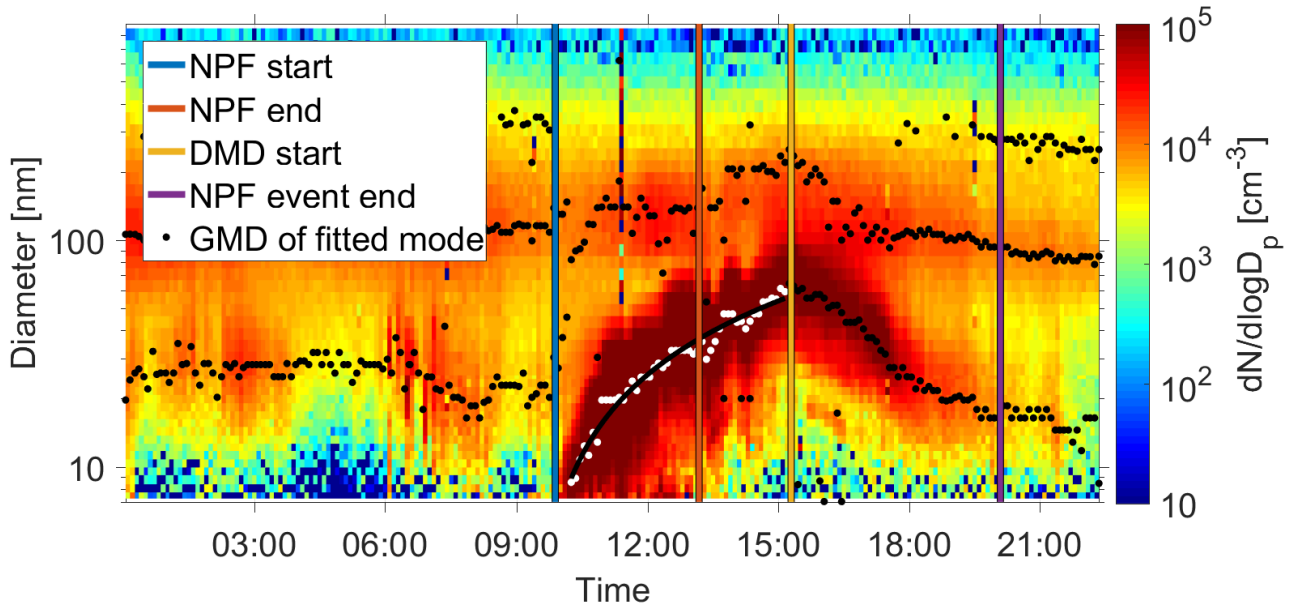
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15 **Figure 1: Particle number size distribution measured by DMPS showing an NPF event with a decreasing mode mean diameter (after 15:00 LT; UTC+3) in Hada Al Sham February 3, 2014. The figure also illustrates the times describing the progression of the NPF events with colored vertical lines (NPF start – blue, NPF end – orange, DMD start – yellow and NPF event end – purple) and the geometric mean diameters (GMDs) of fitted modes with black circles. The GMDs selected for the calculation of the growth rate are shown using white circles, and the black curve shows the linear fit to these points, (colored vertical lines) and the mode fits calculated by the automatic fitting algorithm (black circles) (Hussein et al., 2005). The mode fit points selected for the calculation of the growth rate are shown in white.**

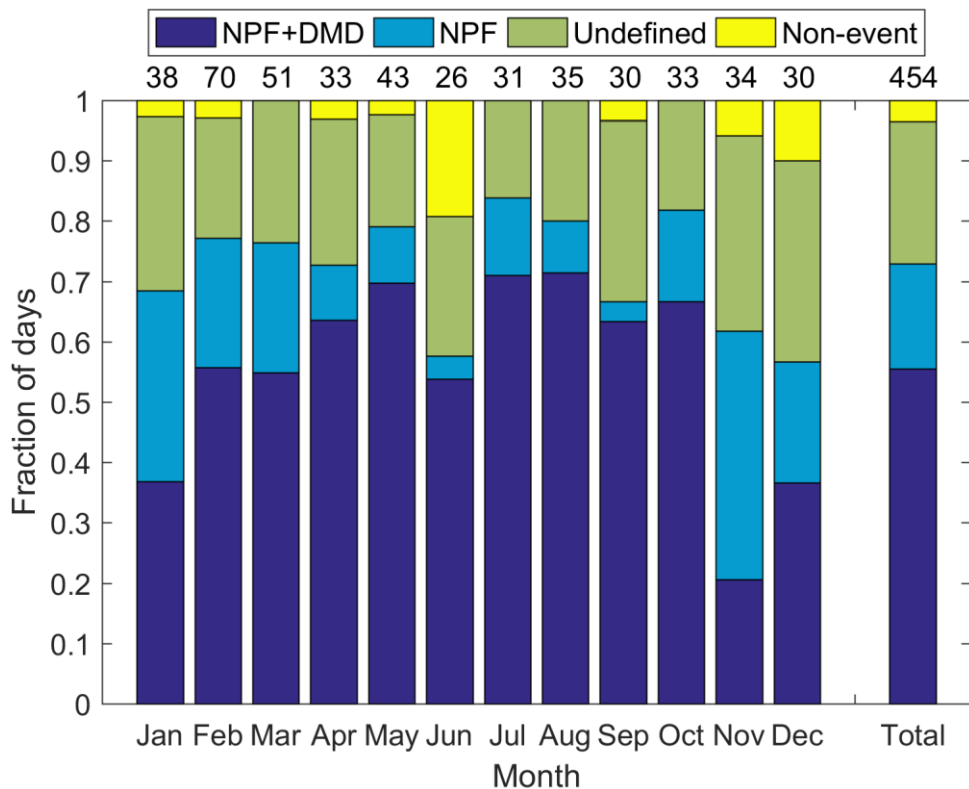


Figure 2: Results of the NPF event classification as fractions of classified days, separately for each month and all of the classified days. The numbers above the bars show the number of classified days. Some of the months contain days from more than a one year, while in some months (June, July, September and December) data was available only for a single year.

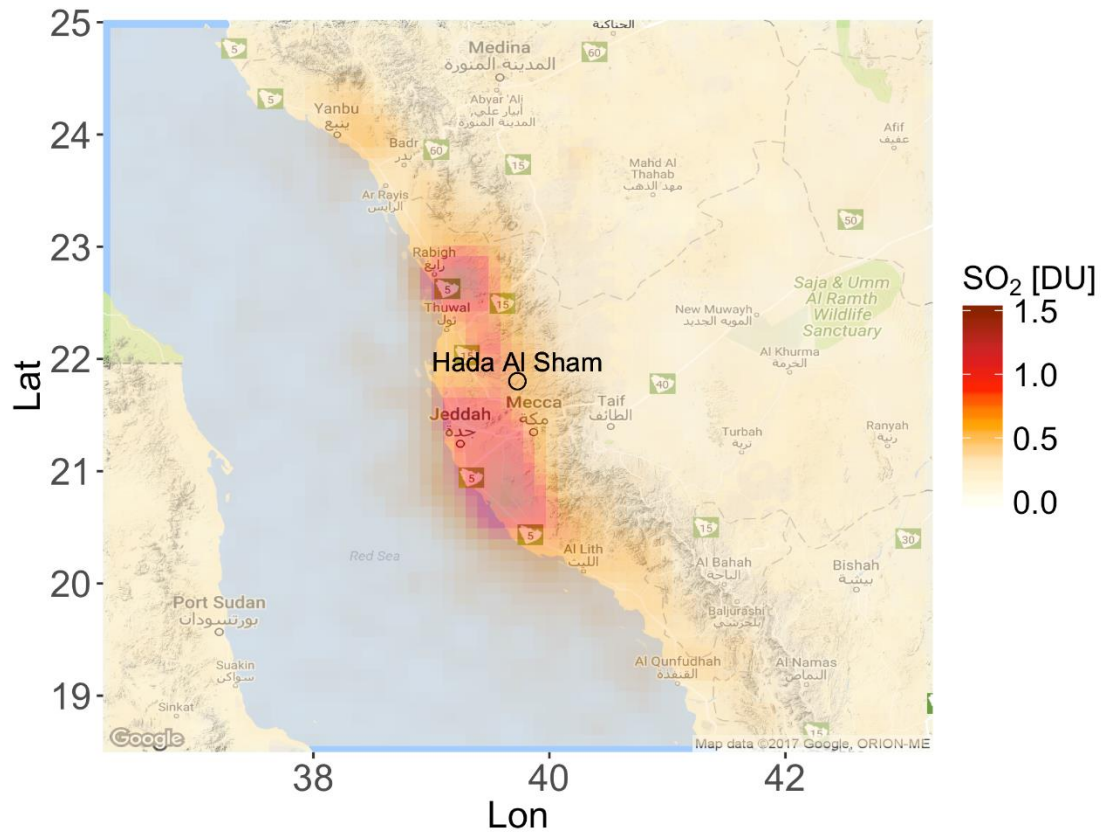


Figure 3: Average concentration of SO₂ from the OMI Level 2 SO₂ Planetary Boundary Layer product (Li et al., 2013) in the surroundings of Hada Al Sham during NPF days. The concentrations are shown in Dobson units (1 DU = 2.69×10^{16} molecules cm⁻²).

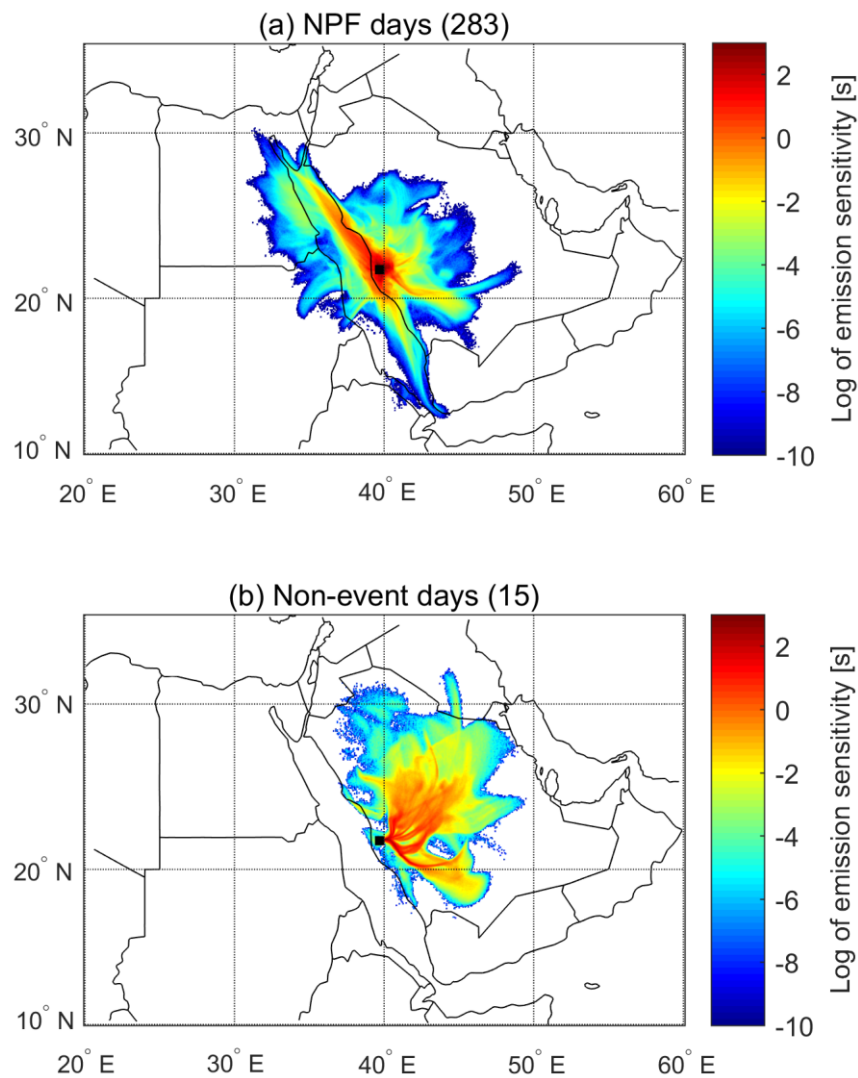


Figure 4: ~~The a~~veraged 24-hour emission sensitivity for air masses arriving at Hada Al Sham at 10:00 LT for (a) NPF days and (b) non-event days. The arrival time of the air masses is chosen to represent the time when new particle formation is typically taking place on NPF days (see Fig. 5).

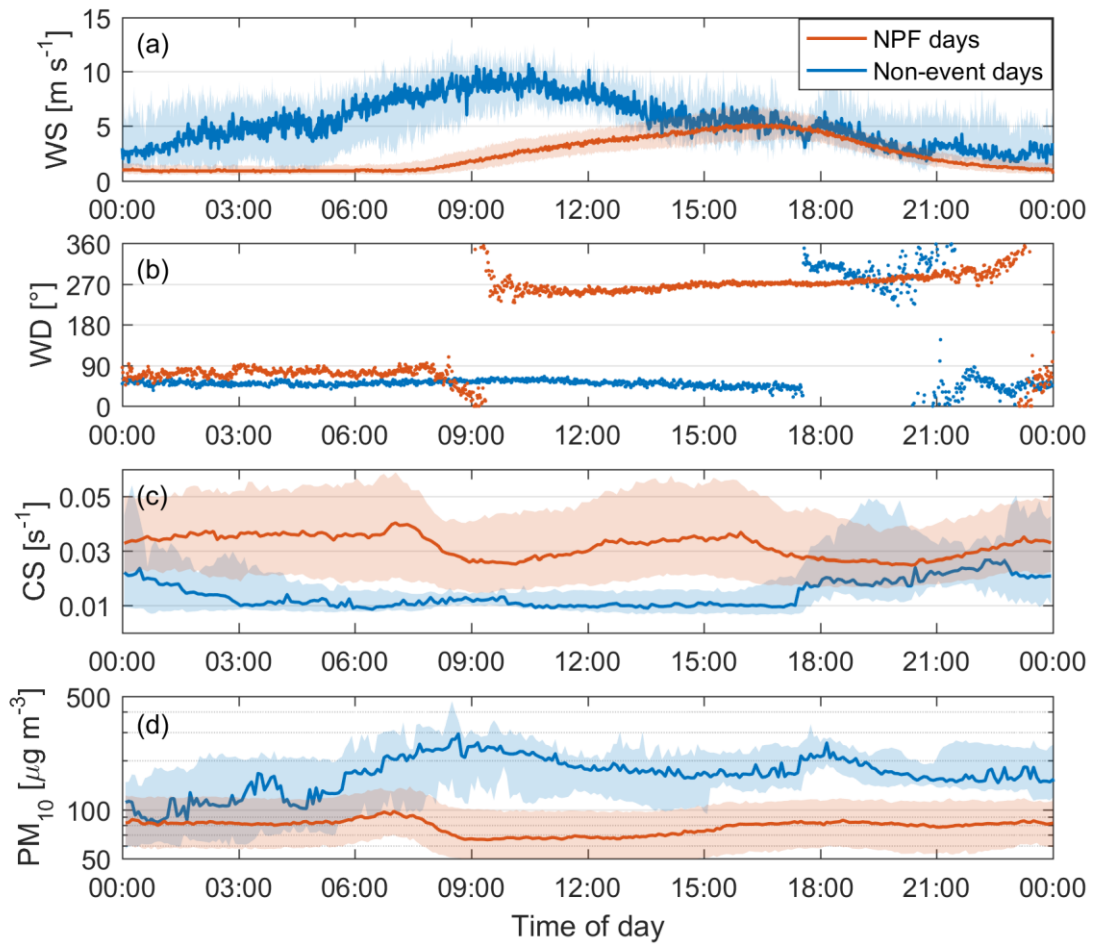


Figure 5: The diurnal variation of (a) wind speed, (b) wind direction, (c) condensation sink and (d) mass of particles smaller than 10 μm in Hada Al Sham during NPF days (red lines) and non-event days (blue lines). The solid lines show the median values and the shaded areas represent the 25th–75th percentile range.

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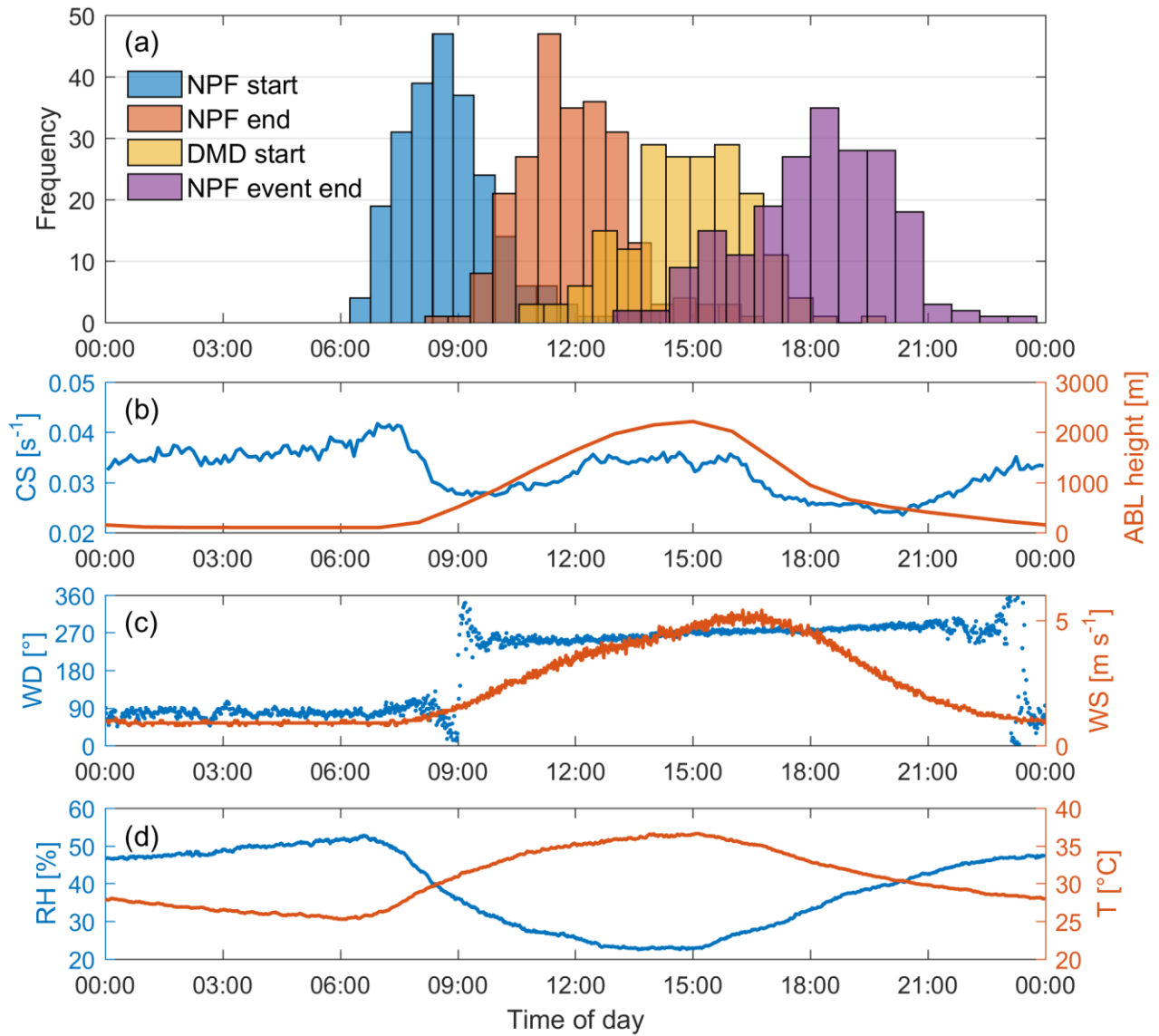


Figure 6: (a) Frequency histograms showing the temporal distribution of the different phases of the NPF events observed in Hada Al Sham (see Fig. 1) together with the median diurnal variation of (b) condensation sink and atmospheric boundary layer height, (c) wind direction and wind speed and (d) relative humidity and temperature.

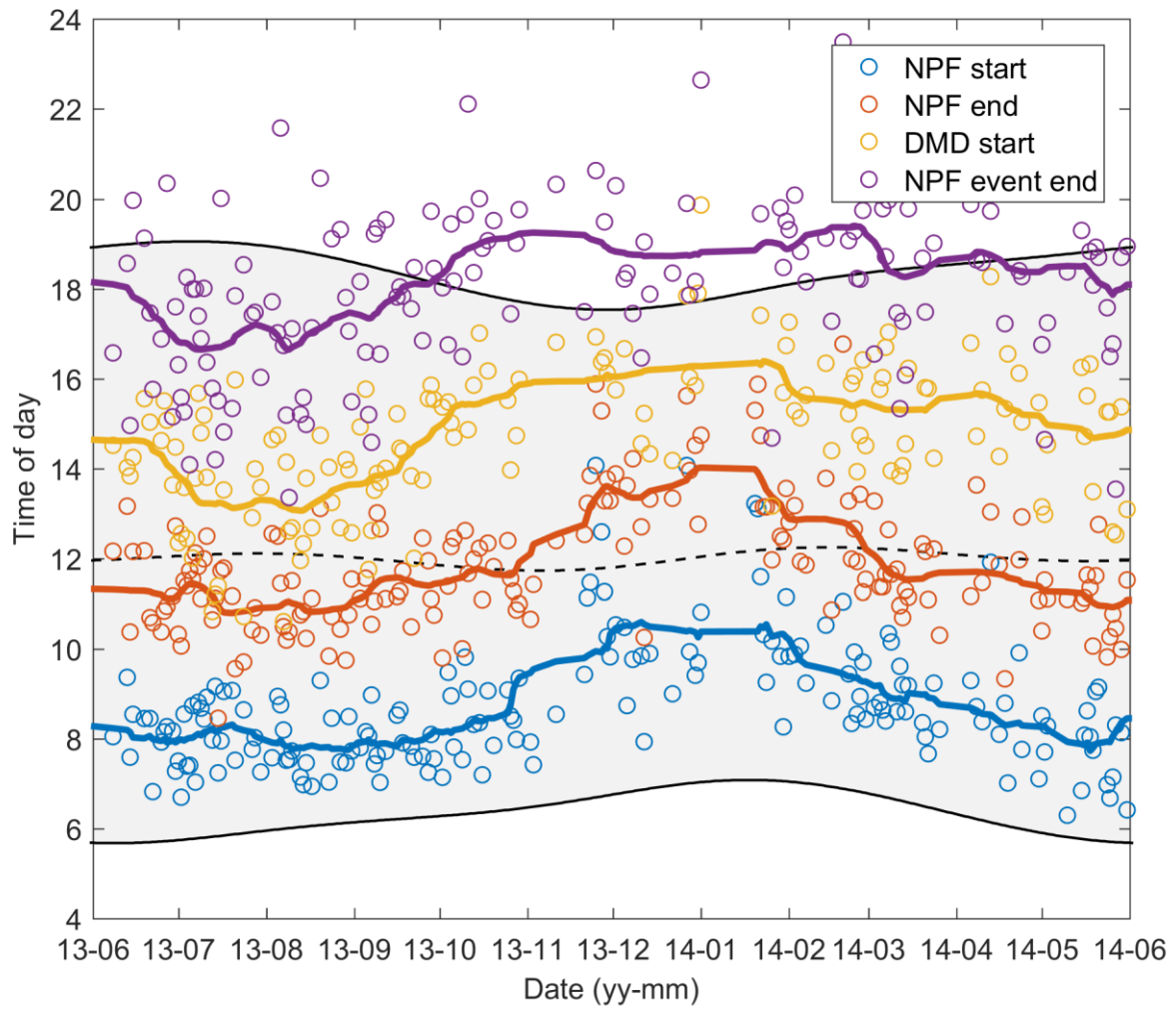
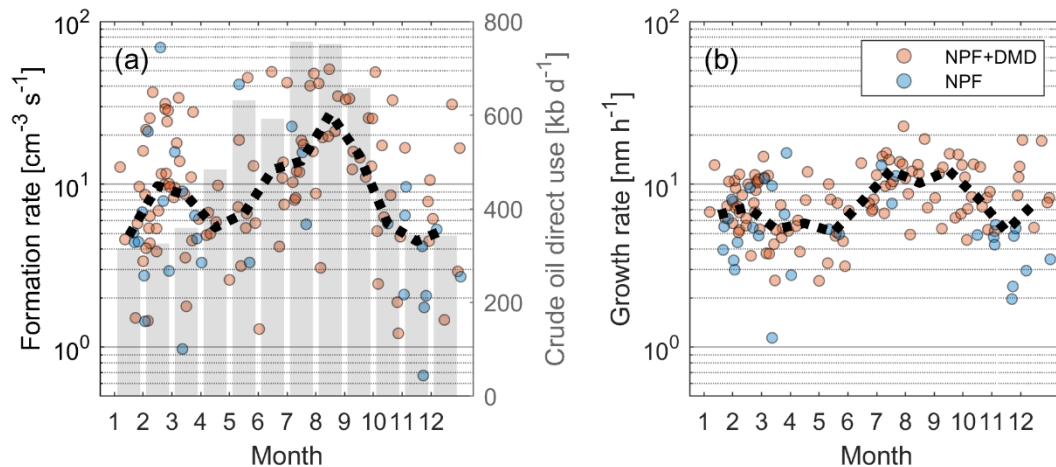
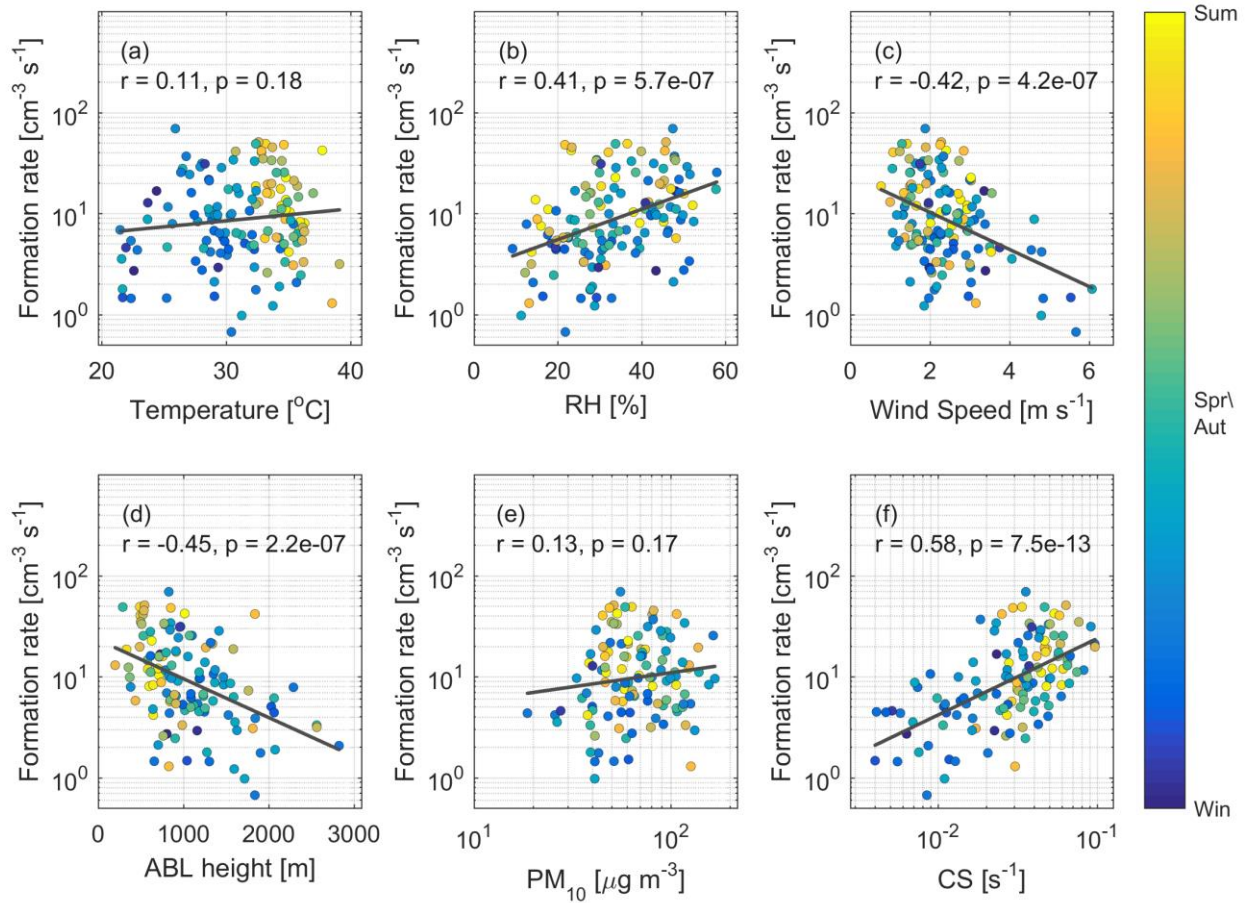


Figure 7: Seasonal variation of the different phases of the NPF events observed in Hada Al Sham. The colored lines show the 20-point moving average for each of the different phases. The solid black lines show the times of sunrise and sunset while the dashed black line shows the time of maximum solar radiation calculated based on the latitude of the measurement site.

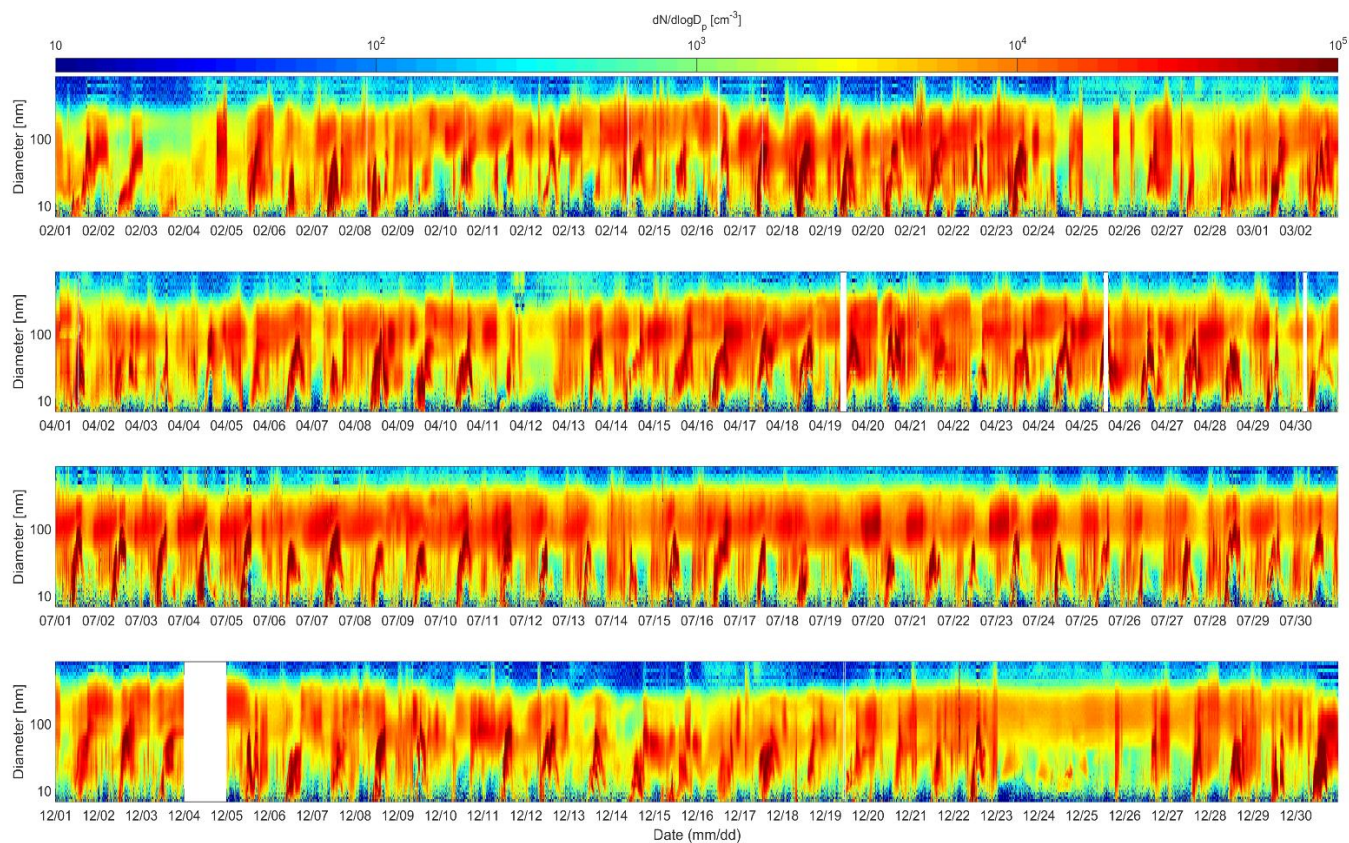


5 **Figure 8: (a) Formation rates ($J_{7\text{nm}}$) and (b) growth rates ($GR_{7-12\text{nm}}$) for the NPF events observed in Hada Al Sham, separately for DMD days (red circles) and non-DMD days (blue circles). The dashed black line represents the monthly medians. The monthly crude oil direct use (<https://www.jodidata.org>) in panel (a) (grey bars, right Y-axis) is calculated as the weighted average of the monthly data from the years 2013–2015. The weighing is based on the number of J values from each year and month.**



5 **Figure 9: Particle formation rate ($J_{7\text{nm}}$) as a function of (a) temperature, (b) relative humidity, (c) wind speed, (d) atmospheric boundary layer height, (e) PM_{10} and (f) the condensation sink. The black lines show the least squares fits to the log-linear (a, b, c, d) or log-log (e and f) data and the r and p -values denote the Pearson's correlation coefficients and their significance levels, respectively. The values on the horizontal axis are event-time averages (from NPF start to NPF end, see Fig. 1). In the case of the CS, the averaging was done for one hour before the NPF start to make sure that the correlation is not influenced by the particles formed in the NPF event itself. The data is colored according to the season, so that summertime is represented by yellow and wintertime by blue colors.**

Appendix A



5 **Figure A1: Time series of particle number size distributions during Feb-2013, Apr-2014, Jul-2013 and Dec-2013 illustrating the high frequency and typical characteristics of NPF, as well as periods of non-event and undefined days e.g. during 23.-27.12.2013. The shown time series were selected so that they would cover different seasons with continuous data availability.**

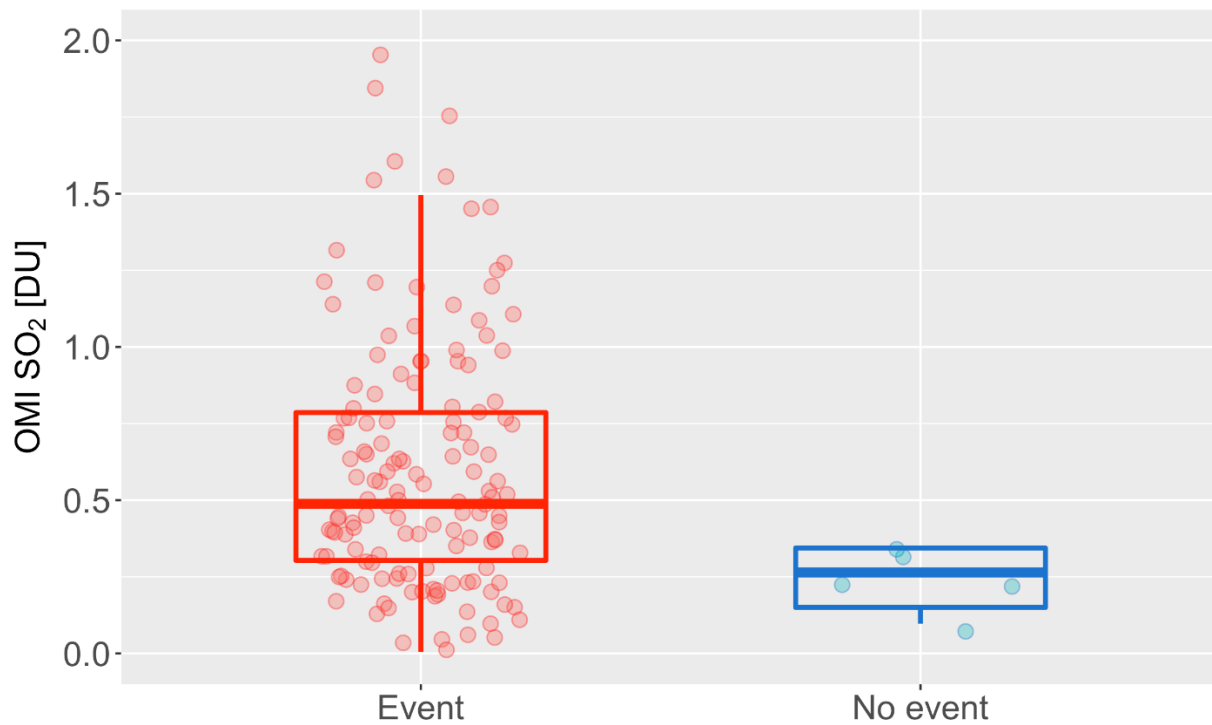


Figure A2: The average SO₂ concentration within the radius of 50 km from Hada Al Sham during NPF days and non-event days from the OMI Level 2 SO₂ Planetary Boundary Layer product (Li et al., 2013). The concentrations are shown in Dobson units (1 DU = 2.69×10^{16} molecules cm⁻²).

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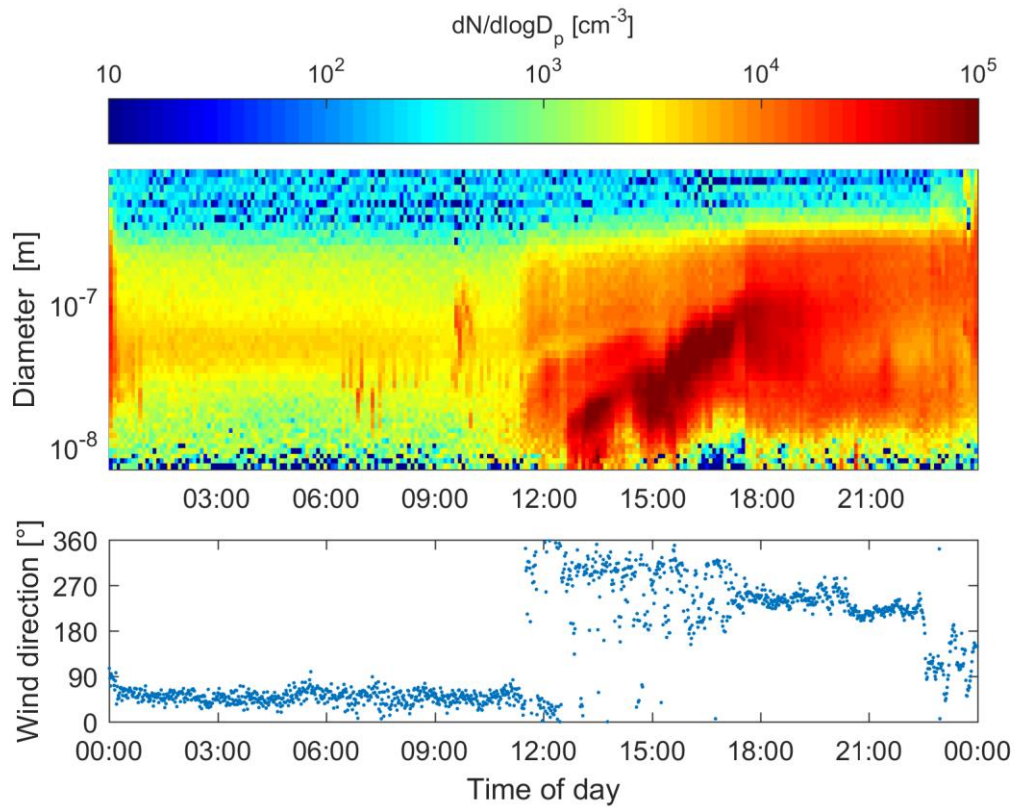


Figure A3: An example illustrating a case where the late starting time of an NPF event is related to a delayed shift in the wind direction in Hada Al Sham, 5 February 2013.

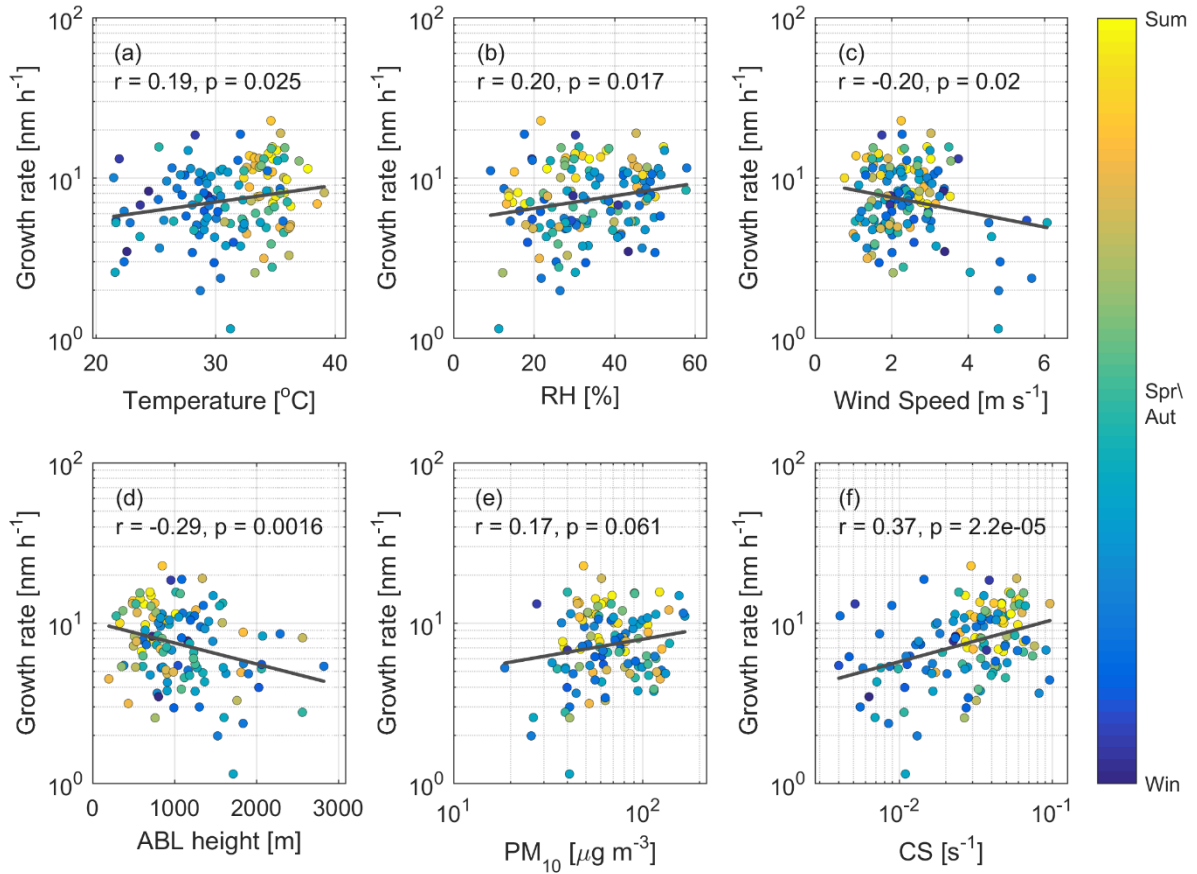


Figure A4: Particle growth rate (GR_{7-12nm}) as a function of (a) temperature, (b) relative humidity, (c) wind speed, (d) atmospheric boundary layer height, (e) PM_{10} and (f) the condensation sink. The black lines show the least squares fits to the log-linear (a, b, c, d) or log-log (e and f) data and the r and p -values denote the Pearson's correlation coefficients and their significance levels, respectively. The values on the horizontal axis are event-time averages (from NPF start to NPF end, see Fig. 1). In the case of the CS, the averaging was done for one hour before the NPF start to make sure that the correlation is not influenced by the particles formed in the NPF event itself. The data is colored according to the season, so that summertime is represented by yellow and wintertime by blue colors.

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