



Retrieval of total column and surface NO₂ from Pandora zenith-sky measurements

Xiaoyi Zhao¹, Debora Griffin¹, Vitali Fioletov¹, Chris McLinden¹, Jonathan Davies¹, Akira Ogyu¹, Sum Chi Lee¹, Alexandru Lupu¹, Michael D. Moran¹, Alexander Cede^{2,3}, Martin Tiefengraber^{3,4}, Moritz Müller^{3,4}

¹Air Quality Research Division, Environment and Climate Change Canada, Toronto, M3H 5T4, Canada.
 ²NASA Goddard Space Flight Center, Greenbelt, MD 20771, USA.
 ³LuftBlick, Kreith 39A, 6162 Mutter, Austria.
 ⁴Department of Atmospheric and Cryospheric Sciences, University of Innsbruck, Innsbruck, Austria.

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Correspondence to: Xiaoyi Zhao (xiaoyi.zhao@canada.ca)

Abstract. Pandora spectrometers can retrieve nitrogen dioxide (NO₂) vertical column densities (VCDs) via two viewing geometries: direct-sun and zenith-sky. The direct-sun NO₂ VCD measurements have high quality (0.1 DU accuracy in clear-sky conditions) and do not rely on any radiative transfer model to calculate air mass factors (AMFs); however, they are not

- 15 available when the sun is obscured by clouds. To perform NO₂ measurements in cloudy conditions, a simple but robust NO₂ retrieval algorithm is developed for Pandora zenith-sky measurements. This algorithm derives empirical zenith-sky NO₂ AMFs from coincident high-quality direct-sun NO₂ observations. Moreover, the retrieved Pandora zenith-sky NO₂ VCD data are converted to surface NO₂ concentrations with a scaling algorithm that uses chemical-transport-model predictions and satellite measurements as inputs. NO₂ VCDs and surface concentrations are retrieved from Pandora zenith-sky measurements
- 20 made in Toronto, Canada, from 2015 to 2017. The retrieved Pandora zenith-sky NO₂ data (VCD and surface concentration) show good agreement with both satellite and in situ measurements. The diurnal and seasonal variations of derived Pandora zenith-sky surface NO₂ data also agree well with in situ measurements (diurnal difference within \pm 2 ppbv). Overall, this work shows that the new Pandora zenith-sky NO₂ products have the potential to be used in various applications such as future satellite validation in moderate cloudy scenes and air quality monitoring.

25 1 Introduction

Nitrogen dioxide (NO₂) is an important air pollutant and plays a critical role in tropospheric photochemistry (e.g., ECCC, 2016; EPA, 2014). It is primarily emitted from combustion processes such as fossil fuel combustion and biomass burning, as well as from lightning. NO₂ is a nitrate aerosol precursor, and it also contributes to acid deposition and eutrophication (ECCC, 2016). Exposure to NO₂ can lead to adverse health effects, such as irritation of the lungs, a decrease in lung

30 function, and an increase in susceptibility to allergens for people with asthma (EEA, 2017; WHO, 2017).





Total vertical column NO_2 can be measured by many ground-based UV-visible remote sensing instruments using direct-sun, zenith-sky, or off-axis spectroscopy techniques (Cede et al., 2006; Drosoglou et al., 2017; Herman et al., 2009; Lee et al., 1994; Noxon, 1975; Piters et al., 2012; Roscoe et al., 2010; Vaughan et al., 1997). These measurements are of high quality and good precision, and have been widely used for atmospheric chemistry studies (e.g. Adams et al., 2012; Hendrick et al.,

- 5 2014) and satellite validations (e.g., Celarier et al., 2008; Drosoglou et al., 2018; Irie et al., 2008; Wenig et al., 2008). Among all these different viewing geometries, direct-sun measurements are of high accuracy, and are not dependent on radiative transfer models (RTMs) to calculate air mass factors (AMFs) (Herman et al., 2009) or on knowledge of other atmospheric constituents. Zenith-sky observations have been widely used for stratospheric ozone and NO₂ observations, particularly under cloudy conditions when direct-sun measurements are unreliable. Off-axis measurements have good
- 10 sensitivity in the boundary layer and could provide tropospheric trace gas profiles and surface concentrations (Frieß et al., 2011; Hendrick et al., 2014; Kramer et al., 2008; Wagner et al., 2011), but they are more sensitive to cloud cover than zenith-sky measurements.

The Pandora sun spectrometer is a new instrument developed to measure vertical column densities (total columns) of trace gases in the atmosphere using sun and sky radiation in the UV-visible part of the spectrum (Herman et al., 2009). One of its

- primary data products is NO₂ total vertical column density (VCD) from the direct-sun viewing mode, where VCD represents the vertically integrated number of molecules per unit area and is reported in units of molec cm⁻² or Dobson Unit (1 DU = 2.6870×10^{16} molec cm⁻²). The Pandora direct-sun NO₂ VCD products have been validated through many field campaigns (Flynn et al., 2014; Lamsal et al., 2017; Martins et al., 2016; Piters et al., 2012; Reed et al., 2015), ground-based comparisons (Herman et al., 2009; Wang et al., 2010), and satellite validations (Ialongo et al., 2016; Lamsal et al., 2014).
- 20 Since their introduction in 2006, Pandora spectrometers have been deployed at more than 50 sites globally. The Pandora no. 103 instrument used in this study has been deployed in Toronto, Canada since 2013 to perform direct-sun measurements (Zhao et al., 2016). Since 2015, the observation schedule of Pandora no. 103 has been modified to perform alternating directsun and zenith-sky measurements. Knepp et al. (2017) assessed Pandora's capability to derive stratospheric NO₂ using zenith-sky viewing geometry (in twilight periods), but their study was limited to slant column densities (SCDs). At this time,
- 25 there are no standard Pandora zenith-sky NO₂ VCD data products available. As one goal of this work, we have focused on developing a new NO₂ retrieval algorithm for zenith-sky measurements to expand Pandora NO₂ measurements into cloudy scenes.

In addition to retrieval of zenith-sky total column NO_2 , another goal of this work is to derive surface NO_2 concentration from total column measurements. Surface NO_2 has been a focus of scientific studies due to its strong correlation with air quality

30 (AQ) and health issues (ECCC, 2016), with NO₂ as one of the three components (along with ozone and PM_{2.5}) used to compute the Air Quality Health Index (AQHI: Stieb et al., 2008) in Canada's AQ public awareness programs. Efforts to link total column NO₂ with its surface concentrations have been made by many researchers (Flynn et al., 2014; Knepp et al., 2015; Kollonige et al., 2017; Lamsal et al., 2008, 2014; McLinden et al., 2014). For example, Knepp et al. (2015) proposed a method to estimate NO₂ surface mixing ratios from Pandora direct-sun total column NO₂ via application of a planetary





boundary-layer (PBL) height correction factor. Kollonige et al. (2017) adapted this method and compared Pandora direct-sun surface NO_2 and OMI surface NO_2 . They concluded that the two main sources of error for the conversion of the total column NO_2 to surface NO_2 are (1) poor weather conditions (e.g., cloud cover and precipitation) and (2) PBL height estimation, both of which affect the NO_2 column-surface relationship and instrument sensitivities to boundary layer NO_2 . Thus, in this work

- 5 we present a simple but robust algorithm for deriving surface NO_2 concentration from Pandora zenith-sky measurements, which has several advantages such as the ability (1) to extend Pandora NO_2 measurements to cloudy conditions and (2) to provide more accurate surface NO_2 concentration estimates that are less sensitive to PBL height. This work also provides reliable total column NO_2 measurements in cloudy conditions and could be used in satellite validations in partially cloudy scenes.
- 10 This paper is organized as follows. Section 2 describes the measured and modelled NO₂ data used in this study. In Section 3, the empirical AMFs for Pandora zenith-sky NO₂ measurements are derived using high-quality Pandora direct-sun total column NO₂ data. These empirical AMFs and the Network for the Detection of Atmospheric Composition Change (NDACC) AMFs (Hendrick et al., 2011; Sarkissian et al., 1995; Van Roozendael et al., 1998; Van Roozendael and Hendrick, 2009; Vaughan et al., 1997) are both applied to Pandora zenith-sky total column NO₂ retrievals to help evaluate
- 15 the performance of the empirical AMFs. Also, the retrieved Pandora zenith-sky total column NO₂ data are evaluated by comparison with satellite measurements. In Section 4, the zenith-sky total column NO₂ data are converted to surface concentration by using a scaling algorithm. The zenith-sky surface NO₂ concentration data are assessed by comparison with in situ measurements. Lastly, in Section 5, several aspects of this zenith-sky surface NO₂ dataset are discussed, which include: diurnal and seasonal variation, and PBL effect, followed by conclusions in Section 6.

20 2 Datasets and models

2.1 Measurements

2.1.1 Pandora direct-sun total column NO₂

The Pandora instrument records spectra between 280 and 530 nm with resolution of 0.6 nm (Herman et al., 2009, 2015; Tzortziou et al., 2012). It uses a temperature-stabilized Czerny-Turner spectrometer, with a 50 µm entrance slit, 1200 groove

 25 mm^{-1} grating, and a 2048×64 back-thinned Hamamatsu charge-coupled device (CCD) detector. The spectra are analysed using a total optical absorption spectroscopy (TOAS) technique (Cede, 2017), in which absorption cross sections for multiple atmospheric absorbers, such as ozone, NO₂, and sulphur dioxide (SO₂), are fitted to the spectra.

The Pandora direct-sun total column NO₂ data are produced using Pandora's standard NO₂ algorithm implemented in the BlickP software (Cede, 2017). The measured direct-sun spectra from 400 to 440 nm are used in the TOAS analysis. A 30 synthetic reference spectrum is produced by averaging multiple measured spectra and corrected for the estimated total optical depth included in it. Cross sections of NO₂ at an effective temperature of 254.5 K (Vandaele et al., 1998), ozone at an





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effective temperature of 225 K (Brion et al., 1993, 1998; Daumont et al., 1992), and a fourth-order polynomial are all fitted. The resulting NO₂ SCDs are then converted to total column VCDs by using direct-sun geometry AMFs. Herman et al. (2009) show that Pandora direct-sun total column NO₂ has a clear-sky precision of 0.01 DU (in slant column) and a nominal accuracy of 0.1 DU (in vertical column). Additional information on Pandora calibrations, operation, and retrieval algorithms can be found in Herman et al. (2009) and Cede (2017).

- The Pandora no. 103 instrument has been deployed in Toronto since September 2013 to perform direct-sun observations (Zhao et al., 2016). The instrument is installed on the roof of the Environment and Climate Change Canada (ECCC) Downsview building (43.7810°N, -79.4680°W) in Toronto. The building is located in a suburban area with multiple roads nearby. Since 2015, the instrument has employed an alternating direct-sun and zenith-sky observation schedule, which
- 10 consists of direct-sun measurements every 90 seconds and zenith-sky measurements every 30 minutes during the sunlit period. About two-and-a-half years (February 2015 to September 2017) of continuous alternating measurements are used in this study.

2.1.2 Pandora zenith-sky total column NO₂

Retrieval of trace gases from Pandora's zenith-sky measurements is not included in the standard BlickP processing software (Cede, 2017). The Pandora zenith-sky spectra for this study are processed using the differential optical absorption spectroscopy (DOAS) technique (Noxon, 1975; Platt, 1994; Platt and Stutz, 2008; Solomon et al., 1987) with the QDOAS software (Danckaert et al., 2015). A single reference spectrum is used, which was obtained from a zenith-sky measurement at local noon from a day that had low total column NO₂. Following the NDACC recommendations (Van Roozendael and Hendrick, 2012), NO₂ differential slant column densities (dSCDs) are retrieved in the 425-490 nm window (to retrieve O₄

20 simultaneously). Cross sections of NO₂ (Vandaele et al., 1998), ozone (Bogumil et al., 2003), H₂O (Rothman et al., 2005), O₄ (Hermans et al., 2003), and Ring (Chance and Spurr, 1997) are all fitted; a fifth-order polynomial and a first-order linear offset are also included in the DOAS analysis.

The output of QDOAS is NO_2 dSCDs, which can be converted to total column NO_2 via the Langley plot method with the use of the NDACC NO_2 AMF look-up table (LUT) (Van Roozendael and Hendrick, 2012). The NDACC AMF LUT is used here

25 only as a reference since it was primarily developed for retrieval of stratospheric NO_2 . Other empirical zenith-sky NO_2 AMFs have been developed and are used to convert NO_2 dSCDs to total columns. Details about these two different AMFs are given in Section 3.1.

2.1.3 OMI SPv3 data

The Ozone Monitoring Instrument (OMI) is a Dutch-Finnish nadir-viewing UV-visible spectrometer aboard the National Aeronautics and Space Administration (NASA)'s Earth Observing System (EOS) Aura satellite that was launched in July 2004. The OMI instrument measures the solar radiation backscattered by the Earth's atmosphere and surface between 270 and 500 nm with resolution of 0.5 nm (Levelt et al., 2006, 2018). OMI has a 780 × 576 CCD detector that measures at 60





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across-track positions simultaneously, and thus, does not require across-track scanning. Due to this approach, the spatial resolution of the CCD pixels varies significantly along the across-track direction: those pixels near the track centre have ground footprint of 13 km \times 24 km (along-track \times across-track), whereas those close to the track edge (e.g. view zenith angle = 56°) have a ground footprint roughly of 23 km \times 126 km (de Graaf et al., 2016). Note that from 2012 onwards the smallest pixels (across-track positions) can no longer be used and are excluded from the analysis (known as the "row anomaly", i.e.

- Levelt et al., 2018). This means the "smallest" pixels available for an OMI comparison are larger than 13 km \times 24 km. The OMI NO₂ data used in this work are the NASA standard product (SP) (Bucsela et al., 2013; Wenig et al., 2008) version 3.0 Level 2 (SPv3.0) (Krotkov et al., 2017). The NO₂ SCDs are derived using the DOAS technique in the 405-465 nm window (Marchenko et al., 2015). The AMFs used in SPv3.0 are calculated by using 1° \times 1.25° (latitude \times longitude)
- 10 resolution *a priori* NO₂ and temperature profiles from the Global Modeling Initiative (GMI) chemistry-transport model with yearly varying emissions (Krotkov et al., 2017).

2.1.4 In situ measurements

The National Air Pollution Surveillance (NAPS) network was established in 1969 to monitor and assess the quality of ambient (outdoor) air in the populated regions of Canada (<u>https://www.canada.ca/en/environment-climate-</u>

- 15 <u>change/services/air-pollution/monitoring-networks-data/national-air-pollution-program.html</u>, accessed 23 November 2018). NAPS provides accurate long-term air quality data (ozone, NO₂, SO₂, carbon monoxide (CO), fine particulate matter, etc.) of a uniform standard across Canada(e.g., Dabek-Zlotorzynska et al., 2011; Reid and Aherne, 2016). The in situ NO₂ data used in this study were collected at the NAPS Toronto North station (located 100 m away from the Pandora instrument). The site is 186 m above sea level, and the height of the air intake is 4 m above the ground.
- 20 The in situ NO₂ concentration is measured using a photolytic NO₂ instrument (Thermo 42i) that is also sensitive to other gaseous inorganic nitrogen compounds (e.g., nitric acid (HNO₃) and peroxyacetyl nitrate (PAN)) (McLinden et al., 2014). Thus in areas where direct NO_x (nitrogen oxides) emission sources are limited and other nitrogen compounds are present, NO₂ may be overestimated (e.g., in rural areas). For the current site, however, this positive bias has been found to be only about 5%, except for very low NO₂ concentrations (<5 ppbv) (Yushan Su, Ontario Ministry of the Environment, Concentration and Parks, percent computing on the current site).</p>
- 25 Conservation and Parks, personal communication, October 2018).

2.2 Numerical models

Predicted NO₂ fields from three atmospheric chemistry models are used in the algorithm described in Section 4.1 to derive surface NO₂ concentration from Pandora zenith-sky total column NO₂ data. Following McLinden et al. (2014), this work uses the Global Environmental Multi-scale Modelling Air quality and CHemistry (GEM-MACH) regional chemical

30 transport model (CTM) and the GEOS-Chem global CTM to simulate total columns and vertical profiles of tropospheric NO_2 and surface NO_2 concentration. The stratospheric NO_2 partial columns are estimated using OMI satellite data and the Pratmo box-model.





2.2.1 GEM-MACH

GEM-MACH is ECCC's regional air quality forecast model. It is run operationally two times per day to predict hourly surface pollutant concentrations over North America for the next 48 hours (Moran et al., 2009; Pavlovic et al., 2016; Pendlebury et al., 2018). The model consists of an online tropospheric chemistry module (Akingunola et al., 2018; Pavlovic
5 et al., 2016) embedded within the ECCC Global Environmental Multi-scale (GEM) numerical weather prediction model (Côté et al., 1998). Physical and chemical processes represented in GEM-MACH include emissions, dispersion, gas- and aqueous-phase chemistry, inorganic heterogeneous chemistry, aerosol dynamics, and wet and dry removal. The model uses gridded hourly emissions fields based on U.S. and Mexican national inventories from the U.S. Environmental Protection Agency (EPA) Air Emissions Modeling Platform and on Canada's national Air Pollutant Emission Inventory (APEI,

10 <u>https://pollution-waste.canada.ca/air-emission-inventory</u>, accessed 23 November 2018) (Zhang et al., 2018). Currently, only NO_x emissions in the PBL are included in the operational model; free-tropospheric NO_x emissions from lightning and inflight aircraft are not considered. In this work, the GEM-MACH hourly NO₂ vertical profiles from 0 to 1.5 km and surface concentrations are retrieved from archived operational forecasts on the native model grid covering North America at 10 km × 10 km horizontal resolution for the period April 2016 to December 2017. The corresponding grid-box closest to the Pandora

15 location was used in this study.

2.2.2 GEOS-Chem

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The GEOS-Chem chemical transport model (Bey et al., 2001) has been used extensively in the retrieval of tropospheric columns, and has been shown to be capable of reasonably simulating the vertical distributions of NO₂ (Lamsal et al., 2008; Martin et al., 2002; McLinden et al., 2014). The model has a detailed representation of tropospheric chemistry, including aerosols and their precursors (Park et al., 2004). In the simulation used in this study, a global lightning NO_x source of 6 Tg N yr⁻¹ (Martin et al., 2002) was imposed. The model was run on a $1/2^{\circ} \times 2/3^{\circ}$ (latitude × longitude) grid in nested mode over North America, and was driven by assimilated meteorology from the Goddard Earth Observing System (GEOS-5). The modelled NO₂ profiles were used to calculate monthly mean NO₂ partial columns in the free troposphere (1.5 to 12 km), as the GEM-MACH model does not include free-tropospheric NO₂ sources (lightning, in-flight aircraft emissions).

25 2.2.3 Pratmo box-model

Pratmo is a stratospheric photochemical box-model (Brohede et al., 2008; Lindenmaier et al., 2011; McLinden et al., 2000). The model has detailed stratospheric chemistry that includes long-lived species (nitrous oxide (N_2O), methane (CH_4), and water vapor (H_2O)) and halogen families (NO_y , Cl_y , and Br_y) that are based on a combination of three-dimensional model output and tracer correlations (Adams et al., 2017). Heterogeneous chemistry of background stratospheric sulfate aerosols is

30 also included. The model is constrained with climatological profiles of ozone and temperature.





Stratospheric NO_2 has a strong diurnal variation; therefore, diurnal corrections must be applied when OMI stratospheric NO_2 measurements (around local noon) are interpolated to Pandora measurement times. Ratios of modelled stratospheric NO_2 columns are calculated at OMI overpass time and Pandora measurement time. These ratios are multiplied by the OMI measured stratospheric NO_2 to produce stratospheric NO_2 columns corresponding to the time of Pandora measurements. Details about the use of the Pratmo box-model and the calculation of stratospheric NO_2 partial columns are provided in

Section 4.1.

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3 Total column NO₂ retrieval

3.1 Zenith-sky air mass factor

The NDACC UV-visible network uses zenith-sky AMFs in its total column NO₂ retrievals. To improve the overall

- 10 homogeneity of the UV-visible NO₂ column measurements, NDACC recommended using the NO₂ AMF LUT (Van Roozendael and Hendrick, 2012). This LUT is based on climatological NO₂ profiles that are composed of (1) 20-60 km NO₂ profiles developed by Lambert et al. (1999, 2000) and (2) 12-20 km NO₂ profiles derived from SAOZ balloon observations (Van Roozendael and Hendrick, 2012). The NO₂ concentration is set to zero below 12 km altitude. The NO₂ AMFs have been calculated using the UVSPEC/DISORT RTM (Hendrick et al., 2006; Wagner et al., 2007). The parameters used in
- 15 building the LUT are: wavelength, ground albedo, altitude of the station, and solar zenith angle (SZA). Aerosol extinction, ozone, and temperature profiles come from an aerosol model (Shettle, 1989), the U.S. Standard Atmosphere, and the TOMS V8 Climatology, respectively.

The NDACC LUT is designed for stratospheric NO_2 retrievals. Note that the absence of tropospheric NO_2 in the NDACC LUT construction will lead to an underestimation of the total column NO_2 in urban areas. For example, from 2015 to 2017,

20 tropospheric NO₂ accounted for $73 \pm 11 \% (1\sigma)$ of the total column amounts in Toronto (OMI SPv3.0 data). To account for this significant tropospheric NO₂ in urban areas, new empirical AMFs were developed in this study and the NDACC AMF LUT is used for comparison purposes only.

Empirical AMFs are calculated for Pandora zenith-sky NO₂ measurements in such a way that they can be used to retrieve zenith-sky total column NO₂ values that match the high-quality Pandora direct-sun total column NO₂ values. Inferring total

- columns from zenith-sky observations through comparisons with accurate direct-sun observations is a common approach for Brewer and Dobson zenith-sky total ozone measurements (Kerr et al., 1988). For example, in the Brewer instrument zenithsky ozone algorithm, weighted zenith-sky light intensities measured at four wavelengths (F) are expressed as a function of the slant path (μ) and total column ozone (Kerr et al., 1981). The nine semi-empirical coefficients used to derive total column ozone from measured F in the equation are estimated from a set of direct-sun and zenith-sky observations made
- 30 nearly simultaneously (Fioletov et al., 2011). Since Pandora zenith-sky spectra can be analyzed to produce NO₂ dSCDs, instead of finding the link between zenith-sky intensity and total column values, deriving empirical zenith-sky AMFs for Pandora zenith-sky measurements is more straightforward.



The relation between VCD and dSCD can be expressed as:

$$VCD = \frac{dSCD + RCD}{AMF} , \qquad (1)$$

where, RCD is the reference column density that shows the slant column amount of the trace gas in the reference spectrum (Section 2.1.2). If we make an assumption that the coincident direct-sun (DS) and zenith-sky (ZS) measurements sampled

5 the same air mass, then the empirical zenith-sky AMFs can be calculated by assuming $VCD_{DS} = VCD_{ZS}$, which gives

$$VCD_{DS}(SZA) = \frac{dSCD_{ZS}(SZA) + RCD_{ZS}}{AMF_{ZS}(SZA)}.$$
 (2)

Next, we can use nearly-coincident VCD_{DS} and $dSCD_{ZS}$ in a multi-non-linear regression to retrieve AMF_{ZS} and RCD_{ZS} together. To ensure the quality of the retrieved AMF_{ZS} , only high quality direct-sun total column NO_2 data are used with $SZA < 75^{\circ}$. Details about the empirical zenith-sky AMF calculation are shown in Appendix A.

- 10 Figure 1 shows a comparison of the empirical zenith-sky AMFs and NDACC AMFs (calculated for the Toronto measurements). Total column NO₂ can then be retrieved using Eqn. (1) and these two sets of AMFs, where the one based on empirical AMFs is referred to as VCD_{Emp} and the one based on NDACC AMFs is referred to as VCD_{NDACC} . The RCD value used in the retrievals is 0.39 ± 0.01 DU, which is retrieved along with AMF_{Emp} (Appendix A). Figure 2 shows the comparisons of the NO₂ columns measured by zenith-sky and direct-sun methods. The regression analyses were performed
- by using the following coincidence criteria: (1) nearest Pandora direct-sun measurement that was within \pm 5 min of Pandora zenith-sky measurement, (2) SZA < 75°, and (3) Pandora direct-sun total column NO₂ data have assured high quality (BlickP L2 data quality flag for nitrogen dioxide = 0). In general, the VCD_{Emp} and VCD_{NDACC} performed as expected. Compared with VCD_{DS}, the VCD_{NDACC} shows a -25% bias, while the VCD_{Emp} only shows a -4% bias (indicated by the red lines on each panel and their slopes). In addition, VCD_{Emp} shows less SZA dependence than VCD_{DS} (see the increased bias for
- 20 measurements made in larger SZA conditions in Figure 2b). These results confirm that, for urban sites, the tropospheric NO₂ profile should be included when calculating empirical zenith-sky AMFs. In the rest of the paper, only the zenith-sky NO₂ retrieved using empirical AMFs will be discussed. Note that the Pandora zenith-sky total column NO₂ data discussed in Sections 3 are a "clear-sky subset" of Pandora zenith-sky measurements. The assessment of Pandora zenith-sky NO₂ measurements in cloudy conditions are provided in Section 4.

25 **3.2** Comparison with satellite measurements

To illustrate the NO_2 variability over Toronto, Figure 3 shows the time series (2015-2017) from Pandora direct-sun, zenithsky, and OMI SPv3.0 total column NO_2 . In general, the NO_2 datasets from the ground-based Pandora instrument and the satellite follow the same pattern. However, the satellite data are likely to miss the peak NO_2 values in the morning since OMI only passes over Toronto once per day around 1:30 p.m. (local time).

30 We also performed regression analyses by using the following coincidence criteria: (1) nearest (in time) measurement that was within \pm 30 min of OMI overpass time, (2) closest OMI ground pixel (having a distance from the ground pixel centre to the location of the Pandora instrument less than 20 km), and (3) cloud fraction ≤ 0.3 (the effective geometric cloud fraction,





as determined by the OMCLDO2 algorithm; Celarier et al., 2016). In this comparison, only high-quality OMI data are used (VcdQualityFlags = 0) (Celarier et al., 2016). Figures 4a and 4b show the scatter plots of OMI vs. Pandora direct-sun and OMI vs. Pandora zenith-sky total column NO₂, respectively. Figures 4c and 4d show similar comparisons but only use OMI NO₂ measured by "small pixels" (i.e., having viewing zenith angle of less than 35°). The better correlation and lower bias for

- 5 zenith-sky versus direct-sun can probably be explained by the fact that the sampled air mass of zenith-sky is closer to what the satellite sees than the air mass sampled by direct-sun measurements. The comparison results indicate that, at the Toronto site, OMI underestimates the total column by about 30 %. This underestimation is qualitatively consistent with the fact that the Pandora location is near the northern edge of peak Toronto NO₂, and the relatively large OMI pixels are also generally sampling areas of less NO₂ in the vicinity. The use of the relatively coarse (1°) GMI model for profiles shapes (Section
- 10 2.1.3) will also lead to a low bias considering the peak NO_x emissions span roughly $0.5^\circ \times 0.5^\circ$. Similar results have been found elsewhere.

Ialongo et al. (2016) reported a similar negative bias using OMI SPv3.0 and Pandora direct-sun total column NO₂ in Helsinki (-32 % bias and R = 0.51), and they suggested this was due to the difference between the OMI pixel and the relatively small Pandora field-of-view. In Reed et al. (2015), Pandora measurements at 11 sites were evaluated; the authors found that the

- 15 best correlation between OMI SPv3.0 and Pandora direct-sun total column NO_2 data is for rural sites. They concluded this could be due to smaller atmospheric variability in the rural region. Other studies such as Goldberg et al. (2017) found an even worse OMI-Pandora comparison between these two data products with striking negative bias at high values and poor correlation (R = 0.3). The authors attributed the poor agreement to the coarse resolution of OMI and its AMFs computed with GMI *a priori* NO₂ profiles. In general, our comparison results show that: (1) the Pandora direct-sun total column NO₂
- 20 data measured in Toronto have a reasonable agreement with OMI, and (2) the Pandora zenith-sky total column NO_2 data show results similar to those for direct-sun total column when compared with OMI SPv3.0.

4 Surface NO₂ concentration retrieval

The performance of the clear-sky Pandora zenith-sky total column NO₂ data has been assessed by using OMI and Pandora direct-sun data as described in Section 3.2. However, the validation of cloudy-scene Pandora zenith-sky total column data is not simple, since near-simultaneous good quality direct-sun or satellite measurements in most cloudy conditions are not available. This cloudy-scene validation can be done by comparison with in situ NO₂ measurements that are not affected by weather. In general, the comparison between total columns and surface concentrations can be done by two approaches: (1) convert Pandora zenith-sky total columns to surface concentrations; and (2) convert in situ surface concentrations to total column values. For example, Spinei et al. (2018) calculated "ground-up" VCDs from in situ surface concentrations by using

30 additional measurements of PBL height or assuming trace gas profiles. In this work, the first approach is employed since the surface NO₂ data products from Pandora remote-sensing measurements have direct applications in areas such as air quality monitoring.





4.1 Column-to-surface conversion algorithm

A simple but robust scaling method is adapted to derive surface NO_2 concentration from Pandora zenith-sky total column NO_2 measurements. Following Lamsal et al. (2008) and McLinden et al. (2014), the surface NO_2 concentration is estimated using the modelled profile and surface concentration,

5 $C_{pan} = \left(V_{pan} - V_{strat} - V_{ftrop}\right) \times \left(\frac{c}{V_{PBL}}\right)_{G-M},$ (3)

where C_{pan} is the surface NO₂ volume mixing ratio (VMR) to be estimated, *C* is the surface NO₂ VMR from GEM-MACH (or G-M), V_{pan} is the total column NO₂ measured by Pandora, V_{strat} is the stratospheric NO₂ partial column, V_{ftrop} is the NO₂ partial column in the free troposphere, and V_{PBL} is the NO₂ partial column in the PBL. This equation assumes the chemical transport models can effectively capture the spatial and temporal behaviour of the concentration-to-partial-column ratio.

- In this work, V_{PBL} (0-1.5 km) is integrated from the GEM-MACH NO₂ profile and V_{ftrop} (1.5-12 km) is integrated from the GEOS-Chem NO₂ profile. Both GEM-MACH and GEOS-Chem have an hourly temporal resolution. Thus, the integrated V_{PBL} and V_{ftrop} can account for NO₂ diurnal variation. However, V_{strat} is from OMI monthly mean stratospheric NO₂, which does not have diurnal variation. Thus, the Pratmo box-model is used to calculate stratospheric NO₂ diurnal ratios. The OMI stratospheric NO₂ columns are interpolated to morning and evening hours by multiplying by the box-model diurnal ratios.
- 15 Details about the calculation of V_{strat} as well as references are provided in Appendix B. The (C/V_{PBL})_{G-M} ratio in Eqn. 3 is provided by GEM-MACH, and has hourly temporal resolution. This modelled (C/V_{PBL})_{G-M} ratio is referred to here as a conversion ratio. Besides the hourly modelled conversion ratio, a simple monthly look-up table is built using an average of the one-and-a-half years of GEM-MACH model outputs (April 2016 to December 2017) that were available. The look-up table (referred to here as the Pandora surface-concentration look-up table, or PSC-LUT) is
- 20 composed of monthly conversion ratios with hourly resolution as shown in Figure 5. For example, assuming that a Pandora NO₂ total column measurement is made on a day in December at 15:00 LST, then the corresponding conversion ratio from the PSC-LUT is 28 ppbv DU^{-1} (see the black arrow). Our results in Figure 5 show that the conversion ratio changes throughout the day as well as with season: 0.1 DU (partial column NO₂ in the PBL) corresponds to 5-8 pptv of surface NO₂ in the morning (8:00 LST), 2-3 pptv around local noon (13:00 LST), and 2-4 pptv in the evening (18:00 LST). In general, the
- variation of conversion ratios demonstrates that the surface NO₂ concentration is controlled not only by PBL height, but also by both boundary-layer dynamics and photochemistry. The surface NO₂ derived using the hourly modelled (C/V_{PBL})_{G-M} ratio is referred to here as $C_{pan-model}$, while the surface NO₂ derived using the monthly mean PSC-LUT is referred to here as $C_{pan-model}$. In general, $C_{pan-model}$ is a data product that depends on daily model outputs, but $C_{pan-LUT}$ only needs the pre-calculated PSC-LUT and is thus less dependent on the model. Details of these two different surface NO₂ data products are discussed in
- 30 the next section.





4.2 Comparison with measurements and model

Figure 6 shows the evaluation of modelled and Pandora zenith-sky surface NO₂ concentrations, both using in situ NO₂ measurements as the reference. The Pandora data have been filtered for heavy clouds (details are given in Section 4.3). The GEM-MACH modelled surface concentrations in Toronto reproduce the in situ measurements very well with the comparison showing high correlation (R = 0.78) and moderate positive bias (37 %, Figure 6a). The Pandora zenith-sky surface NO₂ data, $C_{pan-model}$, shows almost the same correlation (R = 0.77), with only -7 % bias (Figure 6b). The better performance of $C_{pan-model}$ is expected since the conversion method for Pandora zenith-sky measurements relies on the GEM-MACH modelled NO₂ profile (see Eqn. 3); in other words, the Pandora zenith-sky surface NO₂ has at least one more piece of information (i.e., NO₂ total column) than GEM-MACH surface NO₂ concentrations. The $C_{pan-LUT}$ shows a similar correlation coefficient (R = 0.73)

- 10 and has improved bias (-3 %, Figure 6c). This result (slightly lower correlation) is also reasonable and acceptable since $C_{pan-LUT}$ is derived with the monthly PSC-LUT, which has less accurate information than the hourly modelled data. Besides the improved bias, Pandora zenith-sky surface NO₂ concentrations, $C_{pan-model}$ and $C_{pan-LUT}$ (Figures 6e and 6f) also have better frequency distributions than the GEM-MACH (Figure 6d). Figure 6d shows that the NO₂ surface concentrations peaks (ambient background concentrations) from model and in situ data are misaligned. This indicates that the GEM-MACH
- NO₂ background surface concentrations have a 1ppbv low bias at this site. In contrast, the zenith-sky surface NO₂ at peakfrequency matches the in situ data (Figures 6e and 6f), indicating that the low bias of the background surface NO₂ value has been corrected with this additional information from Pandora zenith-sky total column measurements. In addition, in high NO₂ concentration conditions (> 20 ppbv), the zenith-sky surface NO₂ also shows better agreement with the in situ NO₂ than do the modelled data. The mean of the top 10 % of the in situ data is 26 ± 1 ppbv (uncertainty of the mean), whereas the
- 20 corresponding values for GEM-MACH, $C_{pan-model}$, and $C_{pan-LUT}$ are 39 ± 1 ppbv, 26 ± 1 ppbv, and 27 ± 1 ppbv, respectively. The total column-to-surface concentration conversion algorithm has also been applied to the Pandora direct-sun total column NO₂ (see Figure 7). Figure 7b shows that the direct-sun surface NO₂ data have a similar agreement with the in situ data (-8 % bias and R = 0.80) as the zenith-sky surface NO₂. In high NO₂ concentration conditions, direct-sun data have a similarly good agreement with the in situ measurements. For this direct-sun based dataset, the mean of the top 10 % of the in situ data
- is 27 ± 1 ppbv, whereas the corresponding values for GEM-MACH, $C_{pan-model}$, and $C_{pan-LUT}$ are 40 ± 1 ppbv, 27 ± 1 ppbv, and 27 ± 1 ppbv, respectively

Thus, in general, both Pandora zenith-sky and direct-sun surface NO_2 datasets can be used reliably to obtain surface concentrations.

4.3 Measurements in different sky conditions

30 Although zenith-sky observations are less sensitive to cloud conditions than direct-sun observations, we still need to be cautious about the derived zenith-sky surface NO_2 in heavy cloud conditions. Due to enhanced scattering, heavy clouds could lead to a significant overestimation of surface NO_2 derived from zenith-sky measurements. A cloud filtering method





based on retrieved O_4 dSCDs is used to identify these conditions. High retrieved O_4 values correspond to long optical path lengths and therefore it is expected that corresponding NO_2 values are overestimated as discussed in Appendix C.

- The effectiveness of the zenith-sky NO₂ in cloudy scenes is demonstrated by the time series plots (Figure 8) of in situ and Pandora direct-sun and zenith-sky data (in their original temporal resolutions). Under clear-sky conditions (for example, April 8-14), both Pandora direct-sun and zenith-sky-based surface concentrations correlate well with the in situ measurements. Under moderately cloudy conditions, when Pandora direct-sun observations cannot provide high-quality data, Pandora zenith-sky observation still can yield good measurements that compare well with in situ data (for example, April 26-29). Under heavy cloud conditions, however, which are identified by enhanced O₄ (Appendix C), Pandora zenith-sky-derived surface NO₂ yielded higher than in situ measurements (for example, April 4 and 6, see the green squares). This feature is due
- 10 to the enhanced multi-scattering in heavy cloud conditions, which leads to enhanced NO₂ absorption in the measured spectra. Sensitivity tests (Appendix C) show that only 10 % of all zenith-sky measurements are strongly affected by this enhanced absorption, indicating the zenith-sky NO₂ algorithm is applicable to most measurements made in thin and moderate cloud conditions (Toronto has about 44 % of daylight hours with clear-sky conditions per year). The relative strength of direct-sun measured by a collocated Total Sky Imager (model TSI-880) is plotted on top of each panel in Figure 8 as an additional
- 15 indicator of sky conditions. The relative strength of direct-sun is from the integration of blocking-strip luminance. In general, when the relative strength of direct-sun is high (> 60), good quality direct-sun and zenith-sky NO₂ data can both be produced. However, when sun strength is moderate (30-60), only zenith-sky NO₂ data are reliable. When sun strength is low (< 30), zenith-sky NO₂ has increased bias and needs to be filtered out.

5 Discussion

20 This study evaluated the performance of Pandora zenith-sky measurements with Pandora direct-sun measurements, satellite measurements, and in situ measurements. In general, the quality of zenith-sky data is affected by three main factors: (1) quality of empirical zenith-sky AMFs; (2) cloud conditions (heavy clouds or moderate/thin clouds); and (3) quality of modelled NO₂ profile (this factor only applies to Pandora surface NO₂ data). The quality of empirical zenith-sky AMFs and the cloud effect have been addressed in Appendices A and C, respectively. The third factor is discussed in Sections 5.1 and 25 5.2.

5.1 Diurnal and seasonal variation

From the Pandora zenith-sky and direct-sun measurements, and modelled NO_2 profiles, surface NO_2 concentrations were obtained that agree well with in situ measurements collected at the same location. The Pandora surface NO_2 data were also analyzed in more detail with a focus on temporal variations. Figure 9 shows the averaged surface NO_2 diurnal variations of

30 four different datasets. The in situ instrument produces continuous measurements 24 hours per day, whereas Pandora only has measurements when sunlight is available. The diurnal variation of surface NO_2 concentration is controlled by dynamics





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(e.g., vertical mixing, wind direction), photochemistry, and local emissions. Thus, the diurnal variations are calculated using only the hours when in situ, direct-sun, and zenith-sky data are all available.

Figure 9 shows that all four datasets/curves captured the enhanced morning surface NO_2 and the decreasing trend afterwards. However, the model has a positive offset (6-9 ppbv) in the morning (due in part to the use of older emissions inventories: Moran et al., 2018) and a negative offset (1-3 ppbv) in the evening relative to the in situ data. Compared to the modelled data, the Pandora direct-sun and zenith-sky data show improvements in the morning, but almost no changes for the evening.

- This feature is investigated and found to be correlated with the GEM-MACH modelled PBL height (details in Section 5.2). The diurnal variation is also examined by grouping the data by seasons. Figure 10 shows that the surface NO_2 concentrations in winter (December, January, and February) are higher than the corresponding values in summer (June, July, and August).
- 10 This difference is mainly due to short sunlit periods and less solar radiation (e.g., increased lifetime of NO₂ and decreased PBL height) in winter. The model has better agreement with the in situ data in summer than in the colder seasons. The best performance of the model is found around local noon, and this feature is not dependent on seasons. Figure 10 also shows that the quality of Pandora zenith-sky and direct-sun surface NO₂ estimates is affected by the quality of GEM-MACH modelled data. For example, Figure 10c shows that in autumn (September, October, and November), GEM-MACH has the largest
- offset in the morning. This error is thus propagated to the Pandora direct-sun surface data, and leads to a larger offset in the 15 morning (than any other season). On the other hand, when GEM-MACH shows a better agreement with in situ measurements (e.g., in spring and summer), Pandora zenith-sky and direct-sun estimates also show better agreement with in situ observations. In general, both Pandora direct-sun and zenith-sky surface NO₂ data show good agreement with in situ measurements in all seasons; the hourly mean values of Pandora surface NO₂ are all well within the 1σ envelope of the in 20 situ measurements.

5.2 Planetary boundary-layer effect

The larger morning offset in modelled surface NO₂ may indicate that the GEM-MACH modelled PBL heights are biased in the morning when the boundary layer is shallow. Figure 11 (left column) shows the modelled PBL height plotted as a function of the difference between modelled and in situ surface NO₂. Figure 11a shows that, in general, the difference

between modelled and in situ NO₂ decreases with an increase of PBL height. When the modelled PBL height is less than 100 25 m, the mean difference is 18 ± 12 ppbv (1 σ), while when the modelled PBL height is 1 km, the mean difference is only 2.9 \pm 6.4 ppbv.

Even though the modelled surface concentrations are significantly impacted by the PBL, the modelled conversion ratio (from column to surface concentrations) seems unaffected since the surface NO₂ concentrations derived from Pandora zenith-sky

data (C_{pan-model}) show much less dependence on the PBL height (Figure 11b). When the modelled PBL height is less than 100 30 m, the mean difference is 0.9 ± 8.9 ppbv. When the modelled PBL height is 1 km, the mean difference is slightly improved to 0.1 ± 4.4 ppby. Figures 11c and 11h show similar plots as Figure 11a and 11b, but the dataset has been divided into three time-bins (before 9:00, 11:00 to 13:59, and after 15:00). Figures 11c, 11e, and 11f confirm that whenever the modelled PBL





height is low, the relative difference between the model and in situ data is high. However, in general, most of these shallow PBL height conditions occur in the morning, and thus the modelled surface NO_2 has larger bias compared to in situ data in the morning. Figures 11d, 11f, and 11h show that Pandora zenith-sky surface NO_2 data have similar performance for all these three time-bins, which indicates that the data have less PBL height dependency than the modelled data. In other words, the model is able to capture the ratio between the boundary layer partial column and surface NO_2 , although the PBL height

5 the model is able to capture the ratio between the boundary layer partial column and surface NO₂, although the PBL height may not be correct in the model. When this ratio is applied to both Pandora direct-sun and zenith-sky data, the estimated surface concentrations agree better with the in situ measurements.

10 6 Conclusions

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The Pandora spectrometer was originally designed to retrieve total columns of trace gases such as ozone and NO_2 from direct-sun spectral measurements in the UV-visible spectrum. In this work, a new zenith-sky total column NO_2 retrieval algorithm has been developed. The algorithm is based on empirical AMFs derived from nearly simultaneous direct-sun and zenith-sky measurements. It is demonstrated that this algorithm can retrieve total columns in thin and moderate cloud conditions when direct-sun measurements are not available: only 10 % of the measurements affected by heavy cloud have to be filtered out due to large systematic biases (68 %). The new Pandora zenith-sky total column NO_2 data shows only -4%

- bias compared to the standard Pandora direct-sun data product. In addition, OMI NO₂ SPv3.0 data demonstrate similar biases (-30 % and -29 %, respectively) when compared to direct-sun and zenith-sky Pandora total column NO₂ data. Surface NO₂ concentrations were calculated from Pandora direct-sun and zenith-sky total column NO₂ using column-to-
- 20 surface ratios derived from GEM-MACH regional chemical transport model. The bias between Pandora-based direct-sun and zenith-sky NO₂ surface concentration estimates and in situ measurements is only -8 % and -7 % (with correlation coefficients 0.80 and 0.77), respectively, while the bias between the modelled concentrations and in situ measurements is up to 37 %. The Pandora-based surface NO₂ concentrations also show good diurnal and seasonal variation when compared to the in situ data. High surface NO₂ concentrations in the morning (from 6:00 to 9:00, local standard time) are present in all
- 25 measured and modelled datasets, while, on average, the model overestimates surface NO_2 in the morning by 8.6 ppbv (at 7:00 LST). It appears that the bias in modelled surface NO_2 is related at least in part to an incorrectly diagnosed PBL height. In contrast, the difference between Pandora-based and in situ NO_2 does not show any significant dependence on the PBL height. Thus, to enable a fast and practical Pandora surface NO_2 data production, the use of a pre-calculated conversion ratio PSC-LUT is recommended.
- 30 The new retrieval algorithm for Pandora zenith-sky NO_2 measurements can provide high-quality NO_2 data (both total column and surface concentration) not only in clear-sky conditions, but also in thin and moderate cloud conditions, when direct-sun observations are not available. Long-term Pandora zenith-sky NO_2 data could be used in future satellite validation





for the medium cloudy scenes. Moreover, a column-to-surface conversion look-up table was produced for the Pandora instruments deployed in Toronto; therefore, quick and practical Pandora-based surface NO_2 concentration data can be obtained for air quality monitoring purposes. The variation of conversion ratios in the PSC-LUT demonstrates that the surface NO_2 concentration is controlled not only by the PBL height, but also by both boundary-layer dynamics and photochemistry. This conversion approach can also be used to derive surface concentrations from satellite VCD measurements and thus can be particularly useful for the new generation of geostationary satellite instruments for air quality

10 Data availability. Pandora data are available from the Pandonia network (http://pandonia.net/data/). In situ surface NO₂ data are available from the National Air Pollution Surveillance (NAPS) program (http://maps-cartes.ec.gc.ca/rnspa-naps/data.aspx). OMI NO₂ SPv3.0 data are available from https://disc.gsfc.nasa.gov/. Any additional data may be obtained from Xiaoyi Zhao (xiaoyi.zhao@canada.ca).

monitoring such as the Tropospheric Emissions: Monitoring of Pollution (TEMPO, Zoogman et al., 2014).

15 Author contributions. XZ analyzed the data and prepared the manuscript, with significant conceptual input from DG, VF, and CM, and critical feedbacks from all co-authors. JD, AO, VF, XZ, and SCL operated and managed the Canadian Pandora network. AL, MDM, and DG performed and analyzed the GEM-MACH simulations. AC, MT, and MM operated the Pandonia network and provided critical technical support to the Canadian Pandora measurements and subsequent data analysis.

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References

Adams, C., Strong, K., Zhao, X., Bassford, M. R., Chipperfield, M. P., Daffer, W., Drummond, J. R., Farahani, E. E., Feng,
W., Fraser, A., Goutail, F., Manney, G., McLinden, C. A., Pazmino, A., Rex, M. and Walker, K. A.: Severe 2011 ozone depletion assessed with 11 years of ozone, NO₂, and OClO measurements at 80°N, Geophys. Res. Lett., 39, L05806, doi:10.1029/2011gl050478, 2012.





25

Adams, C., Normand, E. N., McLinden, C. A., Bourassa, A. E., Lloyd, N. D., Degenstein, D. A., Krotkov, N. A., Belmonte Rivas, M., Boersma, K. F. and Eskes, H.: Limb–nadir matching using non-coincident NO₂ observations: proof of concept and the OMI-minus-OSIRIS prototype product, Atmos. Meas. Tech., 9(8), 4103–4122, doi:10.5194/amt-9-4103-2016, 2016.

Adams, C., Bourassa, A. E., McLinden, C. A., Sioris, C. E., Clarmann, T. von, Funke, B., Rieger, L. A. and Degenstein, D.
A.: Effect of volcanic aerosol on stratospheric NO₂ and N₂O₅ from 2002–2014 as measured by Odin-OSIRIS and Envisat-MIPAS, Atmos. Chem. Phys., 17(13), 8063–8080, doi:10.5194/acp-17-8063-2017, 2017.

Akingunola, A., Makar, P. A., Zhang, J., Darlington, A., Li, S.-M., Gordon, M., Moran, M. D. and Zheng, Q.: A chemical transport model study of plume-rise and particle size distribution for the Athabasca oil sands, Atmos. Chem. Phys., 18(12), 8667–8688, doi:10.5194/acp-18-8667-2018, 2018.

10 Bey, I., Jacob, D. J., Yantosca, R. M., Logan, J. A., Field, B. D., Fiore, A. M., Li, Q., Liu, H. Y., Mickley, L. J. and Schultz, M. G.: Global modeling of tropospheric chemistry with assimilated meteorology: Model description and evaluation, J. Geophys. Res., 106(D19), 23073–23095, doi:10.1029/2001JD000807, 2001.

Bogumil, K., Orphal, J., Homann, T., Voigt, S., Spietz, P., Fleischmann, O. C., Vogel, A., Hartmann, M., Kromminga, H., Bovensmann, H., Frerick, J. and Burrows, J. P.: Measurements of molecular absorption spectra with the SCIAMACHY preflight model: instrument characterization and reference data for atmospheric remote-sensing in the 230-2380 nm region, J. Photochem. Photobiol. A, 157, 167–184, doi:10.1016/s1010-6030(03)00062-5, 2003.

Brion, J., Chakir, A., Daumont, D., Malicet, J. and Parisse, C.: High-resolution laboratory absorption cross section of O₃. Temperature effect, Chem. Phys. Lett., 213, 610–612, doi:10.1016/0009-2614(93)89169-I, 1993.

Brion, J., Chakir, A., Charbonnier, J., Daumont, D., Parisse, C. and Malicet, J.: Absorption Spectra Measurements for the Ozone Molecule in the 350–830 nm Region, J. Atmos. Chem., 30, 291–299, doi:10.1023/a:1006036924364, 1998.

Brohede, S., McLinden, C. A., Urban, J., Haley, C. S., Jonsson, A. I. and Murtagh, D.: Odin stratospheric proxy NO_y measurements and climatology, Atmos. Chem. Phys., 8(19), 5731–5754, doi:10.5194/acp-8-5731-2008, 2008.

Bucsela, E. J., Krotkov, N. A., Celarier, E. A., Lamsal, L. N., Swartz, W. H., Bhartia, P. K., Boersma, K. F., Veefkind, J. P., Gleason, J. F. and Pickering, K. E.: A new stratospheric and tropospheric NO₂ retrieval algorithm for nadir-viewing satellite instruments: applications to OMI, Atmos. Meas. Tech., 6(10), 2607–2626, doi:10.5194/amt-6-2607-2013, 2013.

Cede, A.: Manual for Blick Software Suite 1.5, 2017.

Cede, A., Herman, J., Richter, A., Krotkov, N. and Burrows, J.: Measurements of nitrogen dioxide total column amounts using a Brewer double spectrophotometer in direct Sun mode, J. Geophys. Res., 111(D5), doi:10.1029/2005JD006585, 2006.

Celarier, E. A., Brinksma, E. J., Gleason, J. F., Veefkind, J. P., Cede, A., Herman, J. R., Ionov, D., Goutail, F., Pommereau,
J.-P., Lambert, J.-C., Roozendael, M. van, Pinardi, G., Wittrock, F., Schönhardt, A., Richter, A., Ibrahim, O. W., Wagner, T., Bojkov, B., Mount, G., Spinei, E., Chen, C. M., Pongetti, T. J., Sander, S. P., Bucsela, E. J., Wenig, M. O., Swart, D. P. J., Volten, H., Kroon, M. and Levelt, P. F.: Validation of Ozone Monitoring Instrument nitrogen dioxide columns, J. Geophys. Res., doi:10.1029/2007JD008908, 2008.

Celarier, E. A., Lamsal, L. N., Bucsela, E. J., Marchenko, S. V. and Krotkov, N. A.: OMNO2 Version 3.0 Level 2 File 35 Description (Document version 3.1), NASA Goddard Space Flight Center., 2016.

Chance, K. V. and Spurr, R. J. D.: Ring effect studies: Rayleigh scattering, including molecular parameters for rotational Raman scattering, and the Fraunhofer spectrum, Appl. Opt., 36, 5224–5230, doi:10.1364/AO.36.005224, 1997.





Côté, J., Gravel, S., Méthot, A., Patoine, A., Roch, M. and Staniforth, A.: The Operational CMC–MRB Global Environmental Multiscale (GEM) Model. Part I: Design Considerations and Formulation, Mon. Wea. Rev., 126(6), 1373–1395, doi:10.1175/1520-0493(1998)126<1373:TOCMGE>2.0.CO;2, 1998.

Dabek-Zlotorzynska, E., Dann, T. F., Kalyani Martinelango, P., Celo, V., Brook, J. R., Mathieu, D., Ding, L. and Austin, C.
C.: Canadian National Air Pollution Surveillance (NAPS) PM2.5 speciation program: Methodology and PM2.5 chemical composition for the years 2003–2008, Atmos. Environ., 45(3), 673–686, doi:10.1016/j.atmosenv.2010.10.024, 2011.

Danckaert, T., Fayt, C., Van Roozendael, M., de Smedt, I., Letocart, V., Merlaud, A. and Pinardi, G.: QDOAS Software user manual version 2.109, Software user manual, Belgian Institute for Space Aeronomy, Brussels., 2015.

Daumont, D., Brion, J., Charbonnier, J. and Malicet, J.: Ozone UV spectroscopy I: Absorption cross-sections at room temperature, J. Atmos. Chem., 15, 145–155, doi:10.1007/bf00053756, 1992.

Drosoglou, T., Bais, A. F., Zyrichidou, I., Kouremeti, N., Poupkou, A., Liora, N., Giannaros, C., Koukouli, M. E., Balis, D. and Melas, D.: Comparisons of ground-based tropospheric NO₂ MAX-DOAS measurements to satellite observations with the aid of an air quality model over the Thessaloniki area, Greece, Atmos. Chem. Phys., 17(9), 5829–5849, doi:10.5194/acp-17-5829-2017, 2017.

15 Drosoglou, T., Koukouli, M. E., Kouremeti, N., Bais, A. F., Zyrichidou, I., Balis, D., van der A, R. J., Xu, J. and Li, A.: MAX-DOAS NO₂ observations over Guangzhou, China; ground-based and satellite comparisons, Atmos. Meas. Tech., 11(4), 2239–2255, doi:10.5194/amt-11-2239-2018, 2018.

ECCC: Canadian Environmental Sustainability Indicators: Air Quality. Environment and Climate Change Canada, ISBN: 978-0-660-06016-3, 2016.

20 EEA: Air quality in Europe — 2017 report, European Environment Agency, EEA Report No. 13/2017, 2017.

EPA: Air Quality Index-A Guide to Air Quality and Your Health, U.S. Environmental Protection Agency, EPA-454/R-00-005, 2014.

Fioletov, V. E., McLinden, C. A., McElroy, C. T. and Savastiouk, V.: New method for deriving total ozone from Brewer zenith sky observations, J. Geophys. Res., 116, doi:10.1029/2010JD015399, 2011.

25 Flynn, C. M., Pickering, K. E., Crawford, J. H., Lamsal, L., Krotkov, N., Herman, J., Weinheimer, A., Chen, G., Liu, X., Szykman, J., Tsay, S.-C., Loughner, C., Hains, J., Lee, P., Dickerson, R. R., Stehr, J. W. and Brent, L.: Relationship between column-density and surface mixing ratio: Statistical analysis of O₃ and NO₂ data from the July 2011 Maryland DISCOVER-AQ mission, Atmos. Environ., 92, 429–441, doi:10.1016/j.atmosenv.2014.04.041, 2014.

Frieß, U., Sihler, H., Sander, R., Pöhler, D., Yilmaz, S. and Platt, U.: The vertical distribution of BrO and aerosols in the
 Arctic: Measurements by active and passive differential optical absorption spectroscopy, J. Geophys. Res., 116, D00R04, doi:10.1029/2011jd015938, 2011.

Goldberg, D. L., Lamsal, L. N., Loughner, C. P., Swartz, W. H., Lu, Z. and Streets, D. G.: A high-resolution and observationally constrained OMI NO₂ satellite retrieval, Atmos. Chem. Phys., 17(18), 11403–11421, doi:10.5194/acp-17-11403-2017, 2017.

35 de Graaf, M., Sihler, H., Tilstra, L. G. and Stammes, P.: How big is an OMI pixel?, Atmos. Meas. Tech., 9(8), 3607–3618, doi:10.5194/amt-9-3607-2016, 2016.





Hendrick, F., Van Roozendael, M., Kylling, A., Petritoli, A., Rozanov, A., Sanghavi, S., Schofield, R., von Friedeburg, C., Wagner, T., Wittrock, F., Fonteyn, D. and De Mazière, M.: Intercomparison exercise between different radiative transfer models used for the interpretation of ground-based zenith-sky and multi-axis DOAS observations, Atmos. Chem. Phys., 6(1), 93–108, doi:10.5194/acp-6-93-2006, 2006.

5 Hendrick, F., Pommereau, J. P., Goutail, F., Evans, R. D., Ionov, D., Pazmino, A., Kyrö, E., Held, G., Eriksen, P., Dorokhov, V., Gil, M. and Van Roozendael, M.: NDACC/SAOZ UV-visible total ozone measurements: improved retrieval and comparison with correlative ground-based and satellite observations, Atmos. Chem. Phys., 11, 5975–5995, doi:10.5194/acp-11-5975-2011, 2011.

Hendrick, F., Müller, J.-F., Clémer, K., Wang, P., Mazière, M. D., Fayt, C., Gielen, C., Hermans, C., Ma, J. Z., Pinardi, G.,
Stavrakou, T., Vlemmix, T. and Roozendael, M. V.: Four years of ground-based MAX-DOAS observations of HONO and NO₂ in the Beijing area, Atmos. Chem. Phys., 14(2), 765–781, doi:10.5194/acp-14-765-2014, 2014.

Herman, J., Cede, A., Spinei, E., Mount, G., Tzortziou, M. and Abuhassan, N.: NO₂ column amounts from ground-based Pandora and MFDOAS spectrometers using the direct-sun DOAS technique: Intercomparisons and application to OMI validation, J. Geophys. Res., 114(D13), D13307, doi:10.1029/2009JD011848, 2009.

15 Herman, J., Evans, R., Cede, A., Abuhassan, N., Petropavlovskikh, I. and McConville, G.: Comparison of ozone retrievals from the Pandora spectrometer system and Dobson spectrophotometer in Boulder, Colorado, Atmos. Meas. Tech., 8, 3407– 3418, doi:10.5194/amt-8-3407-2015, 2015.

Hermans, C., Vandaele, A. C., Fally, S., Carleer, M., Colin, R., Coquart, B., Jenouvrier, A. and Merienne, M.-F.: Absorption cross-section of the collision-induced bands of oxygen from the UV to the NIR, in Weakly Interacting Molecular Pairs:
20 Unconventional Absorbers of Radiation in the Atmosphere, edited by C. Camy-Peyret and A. A. Vigasin, pp. 193–202, Springer, Germany., 2003.

Ialongo, I., Herman, J., Krotkov, N., Lamsal, L., Boersma, K. F., Hovila, J. and Tamminen, J.: Comparison of OMI NO2 observations and their seasonal and weekly cycles with ground-based measurements in Helsinki, Atmos. Meas. Tech., 9(10), 5203–5212, doi:10.5194/amt-9-5203-2016, 2016.

25 Irie, H., Kanaya, Y., Akimoto, H., Tanimoto, H., Wang, Z., Gleason, J. F. and Bucsela, E. J.: Validation of OMI tropospheric NO₂ column data using MAX-DOAS measurements deep inside the North China Plain in June 2006: Mount Tai Experiment 2006, Atmos. Chem. Phys., 8(22), 6577–6586, doi:10.5194/acp-8-6577-2008, 2008.

Kerr, J. B., McElroy, C. T. and Olafson, R. A.: Measurements of ozone with the Brewer ozone spectrophotometer, in Proceedings of the Quadrennial Ozone Symposium, pp. 74–79, Boulder, U.S.A., 1981.

30 Kerr, J. B., Asbridge, I. A. and Evans, W. F. J.: Intercomparison of total ozone measured by the Brewer and Dobson spectrophotometers at Toronto, J. Geophys. Res., 93, 11129–11140, doi:10.1029/JD093iD09p11129, 1988.

Knepp, T., Pippin, M., Crawford, J., Chen, G., Szykman, J., Long, R., Cowen, L., Cede, A., Abuhassan, N., Herman, J., Delgado, R., Compton, J., Berkoff, T., Fishman, J., Martins, D., Stauffer, R., Thompson, A. M., Weinheimer, A., Knapp, D., Montzka, D., Lenschow, D. and Neil, D.: Estimating surface NO₂ and SO₂ mixing ratios from fast-response total column

35 observations and potential application to geostationary missions, J. Atmos. Chem., 72(3–4), 261–286, doi:10.1007/s10874-013-9257-6, 2015.

Knepp, T. N., Querel, R., Johnston, P., Thomason, L., Flittner, D. and Zawodny, J. M.: Intercomparison of Pandora stratospheric NO2 slant column product with the NDACC-certified M07 spectrometer in Lauder, New Zealand, Atmos. Meas. Tech., 10(11), 4363–4372, doi:10.5194/amt-10-4363-2017, 2017.





10

Koenker, R. and Hallock, K. F.: Quantile Regression, J. Econo. Perspec., 15(4), 143–156, doi:10.1257/jep.15.4.143, 2001.

Kollonige, D. E., Thompson, A. M., Josipovic, M., Tzortziou, M., Beukes, J. P., Burger, R., Martins, D. K., Zyl, P. G. van, Vakkari, V. and Laakso, L.: OMI Satellite and Ground-Based Pandora Observations and Their Application to Surface NO₂ Estimations at Terrestrial and Marine Sites, J. Geophys. Res., 123(2), 1441–1459, doi:10.1002/2017JD026518, 2017.

5 Kramer, L. J., Leigh, R. J., Remedios, J. J. and Monks, P. S.: Comparison of OMI and ground-based in situ and MAX-DOAS measurements of tropospheric nitrogen dioxide in an urban area, J. Geophys. Res., 113(D16), doi:10.1029/2007JD009168, 2008.

Krotkov, N. A., Lamsal, L. N., Celarier, E. A., Swartz, W. H., Marchenko, S. V., Bucsela, E. J., Chan, K. L., Wenig, M. and Zara, M.: The version 3 OMI NO₂ standard product, Atmos. Meas. Tech., 10(9), 3133–3149, doi:10.5194/amt-10-3133-2017, 2017.

Lambert, J.-C., Granville, J., Van Roozendael, M., Sarkissian, A., Goutail, F., Müller, J.-F., Pommereau, J.-P. and Russell III, J. M.: A climatology of NO₂ profile for improved Air Mass Factors for ground-based vertical column measurements, in "Stratospheric Ozone 1999", Air Pollution Research Report 73 (CEC DG XII)., 1999.

Lambert, J.-C., Granville, J., Van Roozendael, M., Müller, J.-F., Goutail, F., Pommereau, J.-P., Sarkissian, A., Johnston, P.
V. and Russell III, J. M.: in Global Behaviour of Atmospheric NO₂ as Derived from the Integrated Use of Satellite, Ground-based Network and Balloon Observations, pp. 201–202, Sapporo, Japan., 2000.

Lamsal, L. N., Martin, R. V., Donkelaar, A. van, Steinbacher, M., Celarier, E. A., Bucsela, E., Dunlea, E. J. and Pinto, J. P.: Ground-level nitrogen dioxide concentrations inferred from the satellite-borne Ozone Monitoring Instrument, J. Geophys. Res., 113(D16), doi:10.1029/2007JD009235, 2008.

20 Lamsal, L. N., Krotkov, N. A., Celarier, E. A., Swartz, W. H., Pickering, K. E., Bucsela, E. J., Gleason, J. F., Martin, R. V., Philip, S., Irie, H., Cede, A., Herman, J., Weinheimer, A., Szykman, J. J. and Knepp, T. N.: Evaluation of OMI operational standard NO₂ column retrievals using in situ and surface-based NO₂ observations, Atmos. Chem. Phys., 14(21), 11587– 11609, doi:10.5194/acp-14-11587-2014, 2014.

Lamsal, L. N., Janz, S. J., Krotkov, N. A., Pickering, K. E., Spurr, R. J. D., Kowalewski, M. G., Loughner, C. P., Crawford,
J. H., Swartz, W. H. and Herman, J. R.: High-resolution NO₂ observations from the Airborne Compact Atmospheric Mapper: Retrieval and validation, J. Geophys. Res., 122(3), 1953–1970, doi:10.1002/2016JD025483, 2017.

Lee, A. M., Roscoe, H. K., Oldham, D. J., Squires, J. A. C., Sarkissian, A., Pommereau, J.-P. and Gardiner, B. G.: Improvements to the accuracy of measurements of NO₂ by zenith-sky visible spectrometers, J. Quant. Spectrosc. Radiat. Transfer, 52(5), 649–657, doi:10.1016/0022-4073(94)90031-0, 1994.

30 Levelt, P. F., Hilsenrath, E., Leppelmeier, G. W., Van den Oord, G. H., Bhartia, P. K., Tamminen, J., De Haan, J. F. and Veefkind, J. P.: Science objectives of the ozone monitoring instrument, IEEE Trans. Geosci. Remote Sens., 44, 1199–1208, doi:10.1109/TGRS.2006.872336, 2006.

Levelt, P. F., Joiner, J., Tamminen, J., Veefkind, J. P., Bhartia, P. K., Stein Zweers, D. C., Duncan, B. N., Streets, D. G., Eskes, H., A, R. van der, McLinden, C., Fioletov, V., Carn, S., Laat, J. de, DeLand, M., Marchenko, S., McPeters, R.,

35 Ziemke, J., Fu, D., Liu, X., Pickering, K., Apituley, A., González Abad, G., Arola, A., Boersma, F., Chan Miller, C., Chance, K., Graaf, M. de, Hakkarainen, J., Hassinen, S., Ialongo, I., Kleipool, Q., Krotkov, N., Li, C., Lamsal, L., Newman, P., Nowlan, C., Suleiman, R., Tilstra, L. G., Torres, O., Wang, H. and Wargan, K.: The Ozone Monitoring Instrument: overview of 14 years in space, Atmos. Chem. Phys., 18(8), 5699–5745, doi:10.5194/acp-18-5699-2018, 2018.





25

Lindenmaier, R., Strong, K., Batchelor, R. L., Bernath, P. F., Chabrillat, S., Chipperfield, M. P., Daffer, W. H., Drummond, J. R., Feng, W., Jonsson, A. I., Kolonjari, F., Manney, G. L., McLinden, C., Ménard, R. and Walker, K. A.: A study of the Arctic NO_v budget above Eureka, Canada, J. Geophys. Res., 116(D23), doi:10.1029/2011JD016207, 2011.

Marchenko, S., Krotkov, N. A., Lamsal, L. N., Celarier, E. A., Swartz, W. H. and Bucsela, E. J.: Revising the slant column
density retrieval of nitrogen dioxide observed by the Ozone Monitoring Instrument, J. Geophys. Res., 120(11), 5670–5692, doi:10.1002/2014JD022913, 2015.

Martin, R. V., Chance, K., Jacob, D. J., Kurosu, T. P., Spurr, R. J. D., Bucsela, E., Gleason, J. F., Palmer, P. I., Bey, I., Fiore, A. M., Li, Q., Yantosca, R. M. and Koelemeijer, R. B. A.: An improved retrieval of tropospheric nitrogen dioxide from GOME, J. Geophys. Res., 107(D20), ACH 9-1-ACH 9-21, doi:10.1029/2001JD001027, 2002.

10 Martins, D. K., Najjar, R. G., Tzortziou, M., Abuhassan, N., Thompson, A. M. and Kollonige, D. E.: Spatial and temporal variability of ground and satellite column measurements of NO₂ and O₃ over the Atlantic Ocean during the Deposition of Atmospheric Nitrogen to Coastal Ecosystems Experiment, J. Geophys. Res., 121(23), 14,175-14,187, doi:10.1002/2016JD024998, 2016.

McLinden, C. A., Olsen, S. C., Hannegan, B., Wild, O., Prather, M. J. and Sundet, J.: Stratospheric ozone in 3-D models: A simple chemistry and the cross-tropopause flux, J. Geophys. Res., 105, 14653–14665, doi:10.1029/2000jd900124, 2000.

McLinden, C. A., Fioletov, V., Boersma, K. F., Kharol, S. K., Krotkov, N., Lamsal, L., Makar, P. A., Martin, R. V., Veefkind, J. P. and Yang, K.: Improved satellite retrievals of NO₂ and SO₂ over the Canadian oil sands and comparisons with surface measurements, Atmos. Chem. Phys., 14(7), 3637–3656, doi:10.5194/acp-14-3637-2014, 2014.

Moran, M. D., Menard, S., Talbot, D., Huang, P., Makar, P. A., Gong, W., Landry, H., Gong, S., Gravel, S., Crevier, L.-P.
and Kallaur, A.: Particulate-matter forecasting with GEM-MACH15, a new Canadian operational air quality forecast model, in 30th NATO/SPS ITM on Air Pollution Modelling and Its Application, pp. 289–293, Springer, San Francisco., 2009.

Moran, M. D., Pavlovic, R. and Anselmo, D.: Regional Air Quality Deterministic Prediction System (RAQDPS): Update from version 019 to version 020. Technical note, Sept., Canadian Centre for Meteorological and Environmental Prediction, Montreal, 43 pp., [online] Available from: http://collaboration.cmc.ec.gc.ca/cmc/CMOI/product_guide/docs/tech_notes/technote_raqdps-v20_20180918_e.pdf, 2018.

Noxon, J. F.: Nitrogen dioxide in the stratosphere and troposphere measured by ground-based absorption spectroscopy, Science, 189, 547–549, 1975.

Park, R. J., Jacob, D. J., Field, B. D., Yantosca, R. M. and Chin, M.: Natural and transboundary pollution influences on sulfate-nitrate-ammonium aerosols in the United States: Implications for policy, J. Geophys. Res., 109(D15), doi:10.1029/2003JD004473, 2004.

- Pavlovic, R., Chen, J., Anderson, K., Moran, M. D., Beaulieu, P.-A., Davignon, D. and Cousineau, S.: The FireWork air quality forecast system with near-real-time biomass burning emissions: Recent developments and evaluation of performance for the 2015 North American wildfire season, J. Air Waste Manag. Assoc., 66(9), 819–841, doi:10.1080/10962247.2016.1158214, 2016.
- 35 Pendlebury, D., Gravel, S., Moran, M. D. and Lupu, A.: Impact of chemical lateral boundary conditions in a regional air quality forecast model on surface ozone predictions during stratospheric intrusions, Atmos. Environ., 174, 148–170, doi:10.1016/j.atmosenv.2017.10.052, 2018.





5

Piters, A. J. M., Boersma, K. F., Kroon, M., Hains, J. C., Van Roozendael, M., Wittrock, F., Abuhassan, N., Adams, C., Akrami, M., Allaart, M. A. F., Apituley, A., Beirle, S., Bergwerff, J. B., Berkhout, A. J. C., Brunner, D., Cede, A., Chong, J., Clémer, K., Fayt, C., Frieß, U., Gast, L. F. L., Gil-Ojeda, M., Goutail, F., Graves, R., Griesfeller, A., Großmann, K., Hemerijckx, G., Hendrick, F., Henzing, B., Herman, J., Hermans, C., Hoexum, M., van der Hoff, G. R., Irie, H., Johnston, P. V., Kanaya, Y., Kim, Y. J., Klein Baltink, H., Kreher, K., de Leeuw, G., Leigh, R., Merlaud, A., Moerman, M. M., Monks, P. S., Mount, G. H., Navarro-Comas, M., Oetjen, H., Pazmino, A., Perez-Camacho, M., Peters, E., du Piesanie, A., Pinardi, G., Puentedura, O., Richter, A., Roscoe, H. K., Schönhardt, A., Schwarzenbach, B., Shaiganfar, R., Sluis, W., Spinei, E., Stolk, A. P., Strong, K., Swart, D. P. J., Takashima, H., Vlemmix, T., Vrekoussis, M., Wagner, T., Whyte, C., Wilson, K. M., Yela, M., Yilmaz, S., Zieger, P. and Zhou, Y.: The Cabauw Intercomparison campaign for Nitrogen Dioxide measuring Instruments (CINDI): design, execution, and early results, Atmos. Meas. Tech., 5, 457–485, doi:10.5194/amt-5-457-2012,

10 Instruments (CINDI): design, execution, and early results, Atmos. Meas. Tech., 5, 457–485, doi:10.5194/amt-5-457-2012, 2012.

Platt, U.: Differential optical absorption spectroscopy (DOAS), in Air Monitoring by Spectroscopic Techniques, pp. 27–84, John Wiley, New York., 1994.

Platt, U. and Stutz, J.: Differential Optical Absorption Spectroscopy: Principles and Applications, Springer, Germany, Berlin., 2008.

Reed, A. J., Thompson, A. M., Kollonige, D. E., Martins, D. K., Tzortziou, M. A., Herman, J. R., Berkoff, T. A., Abuhassan, N. K. and Cede, A.: Effects of local meteorology and aerosols on ozone and nitrogen dioxide retrievals from OMI and pandora spectrometers in Maryland, USA during DISCOVER-AQ 2011, J. Atmos. Chem., 72(3–4), 455–482, doi:10.1007/s10874-013-9254-9, 2015.

20 Reid, H. and Aherne, J.: Staggering reductions in atmospheric nitrogen dioxide across Canada in response to legislated transportation emissions reductions, Atmos. Environ., 146, 252–260, doi:10.1016/j.atmosenv.2016.09.032, 2016.

Roscoe, H. K., Van Roozendael, M., Fayt, C., du Piesanie, A., Abuhassan, N., Adams, C., Akrami, M., Cede, A., Chong, J., Clemer, K., Friess, U., Ojeda, M. G., Goutail, F., Graves, R., Griesfeller, A., Grossmann, K., Hemerijckx, G., Hendrick, F., Herman, J., Hermans, C., Irie, H., Johnston, P. V., Kanaya, Y., Kreher, K., Leigh, R., Merlaud, A., Mount, G. H., Navarro, M., Oetjen, H., Pazmino, A., Perez-Camacho, M., Peters, E., Pinardi, G., Puentedura, O., Richter, A., Schonhardt, A.,

- M., Oetjen, H., Pazmino, A., Perez-Camacho, M., Peters, E., Pinardi, G., Puentedura, O., Richter, A., Schonhardt, A., Shaiganfar, R., Spinei, E., Strong, K., Takashima, H., Vlemmix, T., Vrekoussis, M., Wagner, T., Wittrock, F., Yela, M., Yilmaz, S., Boersma, F., Hains, J., Kroon, M., Piters, A. and Kim, Y. J.: Intercomparison of slant column measurements of NO₂ and O₄ by MAX-DOAS and zenith-sky UV and visible spectrometers, Atmos. Meas. Tech., 3, 1629–1646, doi:10.5194/amt-3-1629-2010, 2010.
- 30 Rothman, L. S., Jacquemart, D., Barbe, A., Chris Benner, D., Birk, M., Brown, L. R., Carleer, M. R., Chackerian, J. C., Chance, K., Coudert, L. H., Dana, V., Devi, V. M., Flaud, J. M., Gamache, R. R., Goldman, A., Hartmann, J. M., Jucks, K. W., Maki, A. G., Mandin, J. Y., Massie, S. T., Orphal, J., Perrin, A., Rinsland, C. P., Smith, M. A. H., Tennyson, J., Tolchenov, R. N., Toth, R. A., Vander Auwera, J., Varanasi, P. and Wagner, G.: The HITRAN 2004 molecular spectroscopic database, J. Quant. Spectrosc. Radiat. Transfer, 96, 139–204, doi:10.1016/j.jqsrt.2004.10.008, 2005.
- 35 Sarkissian, A., Roscoe, H. K., Fish, D., Van Roozendael, M., Gil, M., Chen, H. B., Wang, P., Pommereau, J. P. and Lenoble, J.: Ozone and NO₂ Air-Mass Factors for Zenith-Sky Spectrometers - Intercomparison of Calculations with Different Radiative-Transfer Models, Geophys. Res. Lett., 22, 1113–1116, 1995.

Seber, G. A. F. and Wild, C. J.: Nonlinear Regression, Wiley-Interscience, Hoboken, NJ., 2003.

Shettle, E. P.: Models of aerosols, clouds, and precipitation for atmospheric propagation studies, in AGARD Conference 40 Proceedings No. 454: Atmospheric Propagation in the UV, Visible, IR, and MM-Wave Region and Related Systems Aspects, Neuilly sur Seine, France., 1989.





5

Solomon, S., Schmeltekopf, A. and Sanders, R.: On the interpretation of zenith sky absorption measurements, J. Geophys. Res., 2, 8311–8319, 1987.

Spinei, E., Whitehill, A., Fried, A., Tiefengraber, M., Knepp, T. N., Herndon, S., Herman, J. R., Müller, M., Abuhassan, N., Cede, A., Richter, D., Walega, J., Crawford, J., Szykman, J., Valin, L., Williams, D. J., Long, R., Swap, R. J., Lee, Y., Nowak, N. and Poche, B.: The first evaluation of formaldehyde column observations by improved Pandora spectrometers during the KORUS-AQ field study, Atmos. Meas. Tech., 11(9), 4943–4961, doi:10.5194/amt-11-4943-2018, 2018.

Stieb, D. M., Burnett, R. T., Smith-Doiron, M., Brion, O., Shin, H. H. and Economou, V.: A New Multipollutant, No-Threshold Air Quality Health Index Based on Short-Term Associations Observed in Daily Time-Series Analyses, J. Air Waste Manag. Assoc., 58(3), 435–450, doi:10.3155/1047-3289.58.3.435, 2008.

10 Tzortziou, M., Herman, J. R., Cede, A. and Abuhassan, N.: High precision, absolute total column ozone measurements from the Pandora spectrometer system: Comparisons with data from a Brewer double monochromator and Aura OMI, J. Geophys. Res., 117, doi:10.1029/2012JD017814, 2012.

Van Roozendael, M. and Hendrick, F.: Recommendations for total ozone retrieval from NDACC zenith-sky UV-VIS spectrometers, Belgian Institute for Space Aeronomy, Brussels., 2009.

15 Van Roozendael, M. and Hendrick, F.: Recommendations for NO₂ column retrieval from NDACC zenith-sky UV-VIS spectrometers, Belgian Institute for Space Aeronomy, Brussels., 2012.

Van Roozendael, M., Peeters, P., Roscoe, H. K., De Backer, H., Jones, A. E., Bartlett, L., Vaughan, G., Goutail, F., Pommereau, J.-P. and Kyro, E.: Validation of ground-based visible measurements of total ozone by comparison with Dobson and Brewer spectrophotometers, J. Atmos. Chem., 29, 55–83, doi:10.1023/A:1005815902581, 1998.

20 Vandaele, A. C., Hermans, C., Simon, P. C., Carleer, M., Colin, R., Fally, S., Mérienne, M. F., Jenouvrier, A. and Coquart, B.: Measurements of the NO₂ absorption cross-section from 42 000 cm⁻¹ to 10 000 cm⁻¹ (238 -1000 nm) at 220 K and 294 K, J. Quant. Spectrosc. Radiat. Transfer, 59, 171–184, doi:10.1016/s0022-4073(97)00168-4, 1998.

Vaughan, G., Roscoe, H. K., Bartlett, L. M., OConnor, F. M., Sarkissian, A., Van Roozendael, M., Lambert, J. C., Simon, P. C., Karlsen, K., Hoiskar, B. A. K., Fish, D. J., Jones, R. L., Freshwater, R. A., Pommereau, J. P., Goutail, F., Andersen, S.

25 B., Drew, D. G., Hughes, P. A., Moore, D., Mellqvist, J., Hegels, E., Klupfel, T., Erle, F., Pfeilsticker, K. and Platt, U.: An intercomparison of ground-based UV-visible sensors of ozone and NO₂, J. Geophys. Res., 102, 1411–1422, doi:10.1029/96JD00515, 1997.

Wagner, T., Beirle, S., Brauers, T., Deutschmann, T., Frieß, U., Hak, C., Halla, J. D., Heue, K. P., Junkermann, W., Li, X.,
Platt, U. and Pundt-Gruber, I.: Inversion of tropospheric profiles of aerosol extinction and HCHO and NO₂ mixing ratios from MAX-DOAS observations in Milano during the summer of 2003 and comparison with independent data sets, Atmos. Meas. Tech., 4, 2685–2715, doi:10.5194/amt-4-2685-2011, 2011.

<sup>Wagner, T., Burrows, J. P., Deutschmann, T., Dix, B., von Friedeburg, C., Frieß, U., Hendrick, F., Heue, K. P., Irie, H., Iwabuchi, H., Kanaya, Y., Keller, J., McLinden, C. A., Oetjen, H., Palazzi, E., Petritoli, A., Platt, U., Postylyakov, O.,
Pukite, J., Richter, A., Van Roozendael, M., Rozanov, A., Rozanov, V., Sinreich, R., Sanghavi, S. and Wittrock, F.: Comparison of box-air-mass-factors and radiances for Multiple-Axis Differential Optical Absorption Spectroscopy (MAX-DOAS) geometries calculated from different UV/visible radiative transfer models, Atmos. Chem. Phys., 7, 1809–1833, doi:10.5194/acp-7-1809-2007, 2007.</sup>





Wang, S., Pongetti, T. J., Sander, S. P., Spinei, E., Mount, G. H., Cede, A. and Herman, J.: Direct Sun measurements of NO₂ column abundances from Table Mountain, California: Intercomparison of low- and high-resolution spectrometers, J. Geophys. Res., 115(D13), doi:10.1029/2009JD013503, 2010.

Wenig, M. O., Cede, A. M., Bucsela, E. J., Celarier, E. A., Boersma, K. F., Veefkind, J. P., Brinksma, E. J., Gleason, J. F.
and Herman, J. R.: Validation of OMI tropospheric NO₂ column densities using direct-Sun mode Brewer measurements at NASA Goddard Space Flight Center, J. Geophys. Res., 113(D16), doi:10.1029/2007JD008988, 2008.

WHO: Evolution of WHO air quality guidelines: past, present and future, WHO Regional Office for Europe, Copenhagen, 2017.

Zhang, J., Moran, M. D., Zheng, Q., Makar, P. A., Baratzadeh, P., Marson, G., Liu, P. and Li, S.-M.: Emissions preparation
and analysis for multiscale air quality modeling over the Athabasca Oil Sands Region of Alberta, Canada, Atmos. Chem.
Phys., 18(14), 10459–10481, doi:doi.org/10.5194/acp-18-10459-2018, 2018.

Zhao, X., Fioletov, V., Cede, A., Davies, J. and Strong, K.: Accuracy, precision, and temperature dependence of Pandora total ozone measurements estimated from a comparison with the Brewer triad in Toronto, Atmos. Meas. Tech., 9, 5747–5761, doi:10.5194/amt-9-5747-2016, 2016.

15 Zoogman, P., Jacob, D. J., Chance, K., Liu, X., Lin, M., Fiore, A. and Travis, K.: Monitoring high-ozone events in the US Intermountain West using TEMPO geostationary satellite observations, Atmos. Chem. Phys., 14(12), 6261–6271, doi:10.5194/acp-14-6261-2014, 2014.

Appendices

20 A. Empirical zenith-sky AMF

Before calculating the empirical zenith-sky AMF, the VCD_{DS} and $dSCD_{ZS}$ have both been strictly filtered to ensure any measurements used in this calculation have the highest quality. For VCD_{DS} , data are filtered following Cede (2017) with several factors being considered, such as wavelength shift and residual in spectra fitting, direct-sun AMF, and estimated uncertainties for the vertical column. For $dSCD_{ZS}$, data are filtered using similar criteria as for VCD_{DS} , with adjustments for

25 zenith-sky observations.

The VCD_{DS} and dSCD_{ZS} data are merged and divided into several SZA bins. Each bin covers 5° . A multi-non-linear regression is performed by using the following equation:

$$\begin{bmatrix} VCD_1 \\ \vdots \\ VCD_n \end{bmatrix} = \begin{bmatrix} dSCD_1 & \cdots & 0 \\ \vdots & \ddots & \vdots \\ 0 & \cdots & dSCD_n \end{bmatrix} \begin{bmatrix} b_1 \\ \vdots \\ b_n \end{bmatrix} + RCD \begin{bmatrix} I_1 & \cdots & 0 \\ \vdots & \ddots & \vdots \\ 0 & \cdots & I_n \end{bmatrix} \begin{bmatrix} b_1 \\ \vdots \\ b_n \end{bmatrix}$$
(4)

where, VCD_n is not a single direct-sun VCD data point, but is an m × 1 matrix (m is the total number of measurements in 30 SZA bin number *n*); the VCD_n represents all direct-sun VCDs in a 5° SZA bin, and each element of the m × 1 matrix is a single VCD in that SZA bin. Similarly, $dSCD_n$ is also an m × 1 matrix, with each element representing a single coincident zenith-sky dSCD in SZA bin number *n*. I_n is an m × 1 indicator function, where the elements of I_n are set to 1. The *RCD* and





 b_1 to b_n are the parameters to be retrieved. In short, the design of this regression is based on Eqn. 2 (Section 3.1). The idea is to retrieve zenith-sky AMFs in several SZA bins, and, at the same time, all these regressions in different SZA bins are constrained to share a common predictor (RCD). The regression model can be solved by using an iterative procedure (Seber and Wild, 2003) to yield the estimated coefficients, b_1 to b_n and RCD. The b_n is the reciprocal of zenith-sky AMF in SZA bin n.

5 *n*. This regres

This regression model has been evaluated by using different sizes for the SZA bins. A 5° SZA bin is selected because the SZA bin must be small enough to capture the SZA dependency on zenith-sky AMFs, and, at the same time, it must also be large enough to ensure a sufficient number of measurements in each SZA bin (to perform reliable regressions). In order to deal with the diurnal variation of NO₂ concentration and changing of profile shape (e.g., due to changing of boundary layer

10 heights), the dataset has been divided into morning and evening sets, and discrete AMFs are retrieved for a.m. and p.m. separately (see the blue and red squares with error bars in Figure 1).

Next, these discrete AMF values are used to fit an empirical zenith-sky NO₂ AMF function, which has the expression:

$$AMF = a_1 + (1.02 - a_1)/cos(SZA)$$
 (5).

The fitted empirical zenith-sky AMFs are shown in Figure 1 as blue and red lines (data regression period from February

15 2015 to September 2017). Several sensitivity tests have been performed to assess the quality of the empirical zenith-sky AMFs, including fitting the AMFs with/without a diurnal difference, fitting the AMFs with different empirical functions (e.g., exponential and simple geometry approximation) and fitting the AMFs by seasons. All these different choices of empirical AMFs fitting functions or methods only introduce less than 5 % difference in the retrieved empirical AMFs. Thus, to make the empirical AMFs simple and robust, we selected to fit with a diurnal difference (Eqn. 5).

20 B. Stratospheric NO₂ column

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Several stratospheric NO_2 column values were tested and used in the surface NO_2 concentration algorithm (Eqn. 3). Figure A1a shows the OMI monthly mean (referred to as OMI) and Pratmo box-model stratospheric column NO_2 (Adams et al., 2016; McLinden et al., 2000) (referred to as box). Since the satellite only samples Toronto once per day, the OMI stratospheric NO_2 lacks diurnal variation. To account for the diurnal variation, diurnal ratios of NO_2 VCD have been calculated and applied to OMI monthly mean data. The stratospheric NO_2 columns are calculated using

$$V_{OMI}(t) = \frac{V_{box}(t)}{V_{box}(t_0)} \times V_{OMI}(t_0) , \qquad (6)$$

where, $V_{OMI}(t_0)$ is the OMI measured stratospheric column, t_0 is OMI overpass time, $V_{box}(t_0)$ is the modelled stratospheric column at OMI overpass time, $V_{box}(t)$ is the modelled stratospheric column at time t, and $V_{OMI}(t)$ is the interpolated stratospheric columns are referred to as OMI-box. The grey dots on

30 Figure A1b are OMI-box stratospheric NO_2 columns. The monthly mean of the box model (blue line) and OMI-box (black line) show that the amplitude of OMI-box is larger than the amplitude of the box model.





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To justify why this diurnal variation has to be included, Figure A1c shows the total column NO_2 time series. The diurnal stratospheric NO_2 variation is about 0.1 DU in the summer (see grey dots in Figure A1b) when Pandora measured monthly mean total column is about 0.5 DU (Figure A1c). Thus, neglecting this diurnal variation will lead to diurnal biases in the derived surface NO_2 data (e.g., in the morning, this will lead to the overestimation of the stratospheric NO_2 and thus the underestimation of surface NO_2).

C. Cloud effect and heavy cloud filtration

Direct-sun measurements need an unobscured sun. Even thin clouds could decrease the quality of retrieved NO₂ total columns, especially for low altitude clouds. Unlike direct-sun measurements, zenith-sky observations are made with scattered sunlight and have limited sensitivity to cloud cover. For example, Hendrick et al. (2011) calculated that, for NDACC UV-visible zenith-sky ozone measurements, clouds only contribute 3.3 % to the total random error. This is because a trace gas that is mostly distributed in the stratosphere has the mean scattering layer located at a higher altitude than the cloud layer. However, this assumption may not be valid for NO₂. Depending on the properties of the clouds and the NO₂ profile, the clouds could have non-negligible impacts on zenith-sky NO₂ observations.

A typical method of removing zenith-sky measurements affected by heavy clouds is to eliminate measurements with large

enhancements of O_4 and/or H_2O (Van Roozendael and Hendrick, 2012). In the Pandora zenith-sky NO₂ retrieval, we use the O_4 dSCDs. Since the measured O_4 dSCDs has SZA dependency, all measured O_4 dSCDs are plotted against SZA and a second order quantile regression (Koenker and Hallock, 2001) is applied to select the top few percentile of the measured O_4 dSCDs.

Figure A2 shows examples of selected Pandora zenith-sky NO₂ data and their corresponding O₄ dSCDs values. For example,

- Figure A2a shows the O_4 dSCDs versus SZA and the top 10 percentile of the data with enhanced O_4 are marked in grey. The corresponding Pandora zenith-sky data are plotted against in situ data in Figure A2b, which shows low correlation (R = 0.34) and high bias (68 %). This result indicates that the enhanced scattering due to heavy cloud caused a positive bias in the Pandora zenith-sky NO₂ retrieval. Figures A2c and A2d are similar to Figures A2a and A2b, but for selected Pandora zenith-sky NO₂ data that have O_4 values within the 40th to 50th percentile range. Figure A2d shows that when O_4 is not enhanced,
- 25 the derived zenith-sky NO₂ has good agreement with in situ data (R = 0.8 and bias = -5 %). To summarize how the retrieved O₄ dSCDs can indicate the quality of the Pandora zenith-sky NO₂, the results from the other percentile bins are shown in Figure A3. In general, besides the top 10 percentile of data, the results from all the other bins show good correlation (above 0.6) and low bias. Thus, in this study, the Pandora zenith-sky NO₂ data that have O₄ values within only the top 10 percentile are considered to be affected by heavy clouds and are removed. Some examples of this heavy cloud effect are shown in
- 30 Section 4.3.







Figure 1. Comparison of zenith-sky NO₂ air mass factors. Blue and red squares with error bars (standard error) represent the empirical discrete zenith-sky NO₂ AMFs in each SZA bin for Toronto for the period Feb. 2015 to Sept. 2017. Blue and red lines show the fitted empirical zenith-sky NO₂ AMFs. NDACC AMFs calculated using the NDACC look-up table and assuming no NO₂ in the teoposphere are shown in veloce.

5 in the troposphere are shown in yellow.







Figure 2. Comparisons of NO₂ total columns: (a) zenith-sky total column NO₂ retrieved using empirical AMFs vs. direct-sun total column NO₂, (b) zenith-sky total column NO₂ retrieved using NDACC AMFs vs. direct-sun total column NO₂. On each scatter plot, the red line is the linear fit with intercept set to 0, and the black line is the one-to-one line. The scatter plot is colour-coded by solar zenith angle (SZA).







Figure 3. Annual time series of Pandora direct-sun (DS), Pandora zenith-sky (ZS), and OMI SPv3 total column NO_2 in Toronto from 2015-2017.





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Figure 4. OMI vs. Pandora total column NO_2 . (a) and (c) show OMI vs. Pandora direct-sun NO_2 , and (b) and (d) show OMI vs. Pandora zenith-sky NO_2 . (a) and (b) show all available OMI measurements, while (c) and (d) show OMI data from small pixels only. On each scatter plot, the red line is the linear fit with intercept set to 0 and the black line is the one-to-one line. All scatter plots are colour-coded by the distance from the centre of an OMI ground-pixel to the location of Pandora.







Figure 5. Dependence of the Pandora surface NO_2 concentration look-up table (PSC-LUT) on month of year and hour of day. The PSC-LUT is constructed using the GEM-MACH modelled NO_2 conversion ratios. Solid lines are monthly mean conversion ratios colour-coded by month. The shaded envelopes are the standard error of the mean.





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Figure 6. Modelled (a) and Pandora zenith-sky surface NO_2 (b and c) vs. in situ NO_2 . (b) shows the Pandora zenith-sky (ZS) surface NO_2 data derived using the hourly modelled conversion ratio, (c) shows the data derived using the monthly PSC-LUT. (d) to (f) are histograms corresponding to the data in (a) to (c). On each scatter plot, the red line is the linear fit with intercept set to 0 and the black line is the one-to-one line. The scatter plots are colour-coded by the normalized density of the points.





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Figure 7. Modelled (a) and Pandora direct-sun surface NO_2 (b and c) vs. in situ NO_2 . (b) shows the Pandora direct-sun (DS) surface NO_2 data derived using the hourly modelled conversion ratio, (c) shows the data derived using the monthly PSC-LUT. (d) to (f) are histograms corresponding to the data in (a) to (c). On each scatter plot, the red line is the linear fit with intercept set to 0 and the black line is the one-to-one line. The scatter plots are colour-coded by the normalized density of the points.





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Figure 8. Example of surface NO_2 concentration time series in all conditions. The in situ, Pandora direct-sun (DS), and Pandora zenith-sky (ZS) surface NO_2 concentrations are shown by different coloured dots. The total sky imager relative strength of direct-sun data are plotted as a colour-coded horizontal dot-line on the top part of each panel. For Pandora zenith-sky data, the measurements with enhanced O_4 (heavy cloud indicator) are also labelled by green squares.







Figure 9. Diurnal variation of surface NO_2 concentration. The x-axis is the local standard time (LST). Lines with dot/square symbols represent the hourly mean of corresponding data indicated by the legend. The shaded area represents the 1σ envelope.







Figure 10. Diurnal variation of surface NO_2 concentration by season. The x-axis is the local standard time (LST). Each panel represents data collected in one season (spring, summer, autumn or winter). Solid lines represent mean of corresponding data indicated by the legend. The shaded area represents the 1σ envelope.





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Figure 11. Illustration of planetary boundary-layer (PBL) effect. The y-axis is planetary boundary-layer height in km. The x-axes for the left column are the difference between GEM-MACH and in situ surface NO_2 concentrations; the x-axes for the right column are the difference between Pandora zenith-sky ($C_{pan-model}$) and in situ surface NO_2 concentration. Panels a and b show all available data, panels c and d show the morning data (before 9:00, local standard time), panels e and f show the noon data (from 11:00 to 13:59), and panels g and h show the evening data (after 15:00).







Figure A1. Time series of measured and modelled NO_2 columns: (a) stratospheric columns from the box model (hourly) and OMI (monthly), (b) stratospheric columns from OMI-box (hourly), box (monthly) and OMI-box (monthly), and (c) total columns from Pandora zenith-sky and OMI.





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Figure A2. Illustration of cloud effect and heavy-cloud data filtration: (a) shows measured O_4 differential slant column densities versus solar zenith angle, the gray dots represent the top 0.10^{th} percentile range of O_4 , (b) shows the scatter plot of zenith-sky versus in situ surface NO₂ use data that has O_4 value within 0.10^{th} percentile range (as identified in (a)), (c) is similar to (a) but the gray dots represent the 40.50^{th} percentile range of O_4 , (d) is similar to (b) but use the data that has O_4 value within the 40.50^{th} percentile range. On scatter plot (b) and (c), the blue line is the linear fit with intercept set to 0, the red line is a simple linear fit, and the black line is the one-to-one line. All plots are colour-coded by the normalized density of the points.







Figure A3. Correlation coefficient and bias (slope) between zenith-sky and in situ surface NO_2 data in different O_4 dSCDs percentile bins. (a) shows the correlation coefficients, (b) shows the slopes of linear fit with intercept set to 0.