Referee 2:

We thank the reviewer for his very detailed and thorough review. We are confident that we addressed all points raised in the below comments. Regarding the major concerns: We carefully checked the manuscript regarding the "inconsistency with negative signs "and revised the manuscript accordingly. The submitted manuscript was focused on the dynamical causes/limitations of Is but we followed the reviewer's suggestion to additionally look into the causes/limitations of Is resulting from the chemical regime which is driven by NO_x abundance. Furthermore, we give some reasoning why we think that the air-craft and modeling study, whose inclusion was criticized, are comparable to our results and are worth including. We used data from the LES model from the near surface layer 10 m to 30 m for the detailed analysis and the CBL-integrated value only for comparison to other CBL-integrated LES studies like in fig. 13. The study of Kaser at al. (2015) also covers the frequency range (i.e. scales) given by the ground based measurements although it extends to larger scales as well. More detailed comments in the following:

In general: We propose to add a Supplement to the revised version of the paper. Here we give figures S1 - S27 for this Supplement as an example to provide information on the micrometeorological and chemical situation at the ATTO site on day 22/11/2015.

Regarding the NO_x- dependency:

We compared the dependency of *Is* on NO, NO₂, and NO_x mixing ratios and found no significant relation for the measured data (see Fig. S22 in the Supplement). In addition, we give Fig. S15 – S17 for relations between mean OH and mean NO_x, mean isoprene and mean OH and mean Isoprene and mean NO_x (here we also compare to results from Kim et al. (2016)). The results given in Fig. S15 could be compared to those inferred by Liu at al. (2018) as well.

Comparability to Kaser et al. (2015) and Ouwersloot et al. (2011):

According to the Ergodic theorem, one can compare averages over space with averages over time. The link of spatial and temporal scales has been shown for the atmosphere by Orlanski (1975). The scales in space are therefore not the only relevant scales, and only the ranges of frequencies f or wave numbers $k=2\pi f u'$ (u= Wind velocity or flight velocity) are important. Kaser et al. (2015) give horizontal length scales from about 3 km up to 220 km. With the mean measuring velocity of the aircraft of 100 ms⁻¹ one obtains their frequency range of measured data (OH) between $4.5 \cdot 10^{-4}$ Hz and 0.033 Hz. Note that we obtained data for $1.7 \cdot 10^{-3}$ Hz (respectively $5.4 \cdot 10^{-4}$ Hz) up to 0.2 Hz (ECHO) respectively 0.067 Hz (ATTO). Kaser et al. (2015) extrapolated their spectra for OH by a direct proportionality to the spectra of O_3 (not to the chemical production term of OH, which would include the fluctuations of photolysis frequency and water vapor as well). By this procedure, they extended their spectral presentation to higher frequencies. On average they give $Is \sim 0.13$ with a contribution from the smaller scales (>5.10⁻³ Hz) of 0.06 to 0.08 (in absolute values) which is within the range of our measured surface values. Therefore, the data given by Kaser et al. (2015) directly fit into our Fig. 12. The results from Ouwersloot et al. (2011) are taken from their original LES data from mean vertical profiles for the lowest layers (20 m vertical resolution) to compare to our results from heights above rough surfaces.

"Much of this work is a direct reworking of the results presented in Dlugi et al. (2014)":

In our figures Fig. 1 - 14 we present the results of measurements and analysis – and some intermediate steps. We describe what is given by the analysis of data following the concept that terms of balance equations are determined. We don't mix experimental data with modelled data in the analysis of our data in contrast to other sources given in

literature. We then compare the derived results to those form modelling studies. We are aware that not all quantities that are needed to describe this dynamical – chemical – biological system in a complete way (also in the sense of theoretical physics / mathematics) were measured. Therefore, some of them are estimated from additional studies. As this needs to be done for the ATTO data anyway and it is an important step forward to have it done for two contrasting environments ("high-NO_x low isoprene" and "low-NO_x high isoprene") as direct measurements are still very sparse we think that all steps should be included. The point "...to have it all done in one place..." has also been acknowledged by the reviewer as well.

We think it is important not only to focus on quantities that show a "convincing relationship" as a) correlation does not mean causality and b) to give a framework on how data are organized and if different studies can be compared it may be good to see in which state of the system (convection/no convection or dominance of transport/dominance of chemistry) they were derived. Variables describing the state of the system may also be important to decide whether data from different states can be used to derive general relationships between Is and other quantities. This refers especially to Figs. 13 and 14 (Is versus Damköhler number and buoyant production respectively). We do not quite understand the criticism on the relationship between Is and r. The correlation coefficient has been used in modelling studies to derive Is (in absence of other available data) measured. Though derived from very different environments, the ATTO and ECHO data show the same behavior and none of the measured data is below the line with a slope of 2.5. This means that if all data show this behavior even at a perfect anticorrelation (correlation coefficient -1) Is would reach at maximum a value of -0.4. This is an important information from data analysis only. Furthermore, from Fig. 9 it is clear that models tend to have a higher r_{ii} due to the more Gaussian distribution of the data than obtained from field measurements.

Generally, we compared any empirical relations between *Is* and other quantities. As example we take the concept of shear and buoyant generation (or "production") as illustrated in Fig. 4.24 from R. Stull (2000). Corresponding Richardson numbers are given in our Fig. S5.

We noticed for ECHO that a certain empirical relation exists between the "buoyant production BP" and *Is* for a range of BP above $3 \cdot 10^{-3} m^2 s^{-3}$, which is related to free convection conditions. This is given in Fig. 14 for ECHO 2003. The results for ATTO 2015 only partly follow this relation (Fig. 14b in revised figures). In contrast to the results for ECHO 2013 the higher BP (> $3 \cdot 10^{-3} m^2 s^{-3}$) is related to |-Is| < 0.04 because the correlation coefficients become small (see also revised Fig. 9). All these *Is*- data from ECHO 2003 are related to ogives of the covariance $\overline{OH'ISO'}$ with partly negative and partly positive contributions, which sum up to small values of *Is* only in the time interval of 10 minutes. This is described by Dlugi et al. (2014), e.g., in their Fig. 17.

For ATTO 2015 the I_s values for $BP > 5 \cdot 10^{-3}m^2s^{-3}$ are below $|I_s| = 0.04$. The contribution to I_s from frequencies higher $5 \cdot 10^{-3}Hz$ is small. Predominantly, eddies from the low frequency range contribute to the covariance $c'_i c'_j$ respectively the correlation coefficient r_{ij} . Their contribution to I_s is not sufficiently covered by a ten – minute averaging interval during situations where cloud and surface induced vertical and horizontal convective mixing interact. This is also illustrated by Figures S26 and S27 for I_s as function of surface sensible heat flux H_s . (Note that I_s is always given at the end of the 10 – minute interval and, therefore, is shifted to the right compared to maxima or maxima in H_s .) Therefore, above about $BP = 3.5 \cdot 10^{-3} m^2 s^{-3}$ only one data point shows a larger I_s ($I_s = -0.067$), while the others are influenced by the effect mentioned above. Note that we choose a 10- minute

interval to approach conditions for stationarity (see our remarks on data analysis and covariance calculations). If we would extend the range to lower frequencies, as done for ECHO, stationarity conditions would not be fulfilled.

The relation given in Fig. 13 is again an empirical presentation of the data on Is, which shows that for moderately unstable conditions all data point towards an increasing (with increasing Da) but limited Is. Some points, where " u_* - scaling" is not fulfilled, don't follow such behavior.

Line-by-line concerns:

40 - 43	Rev: Lines 40-43: Please give an approximate range of NO_x for these conditions (instead of "high NO_x " and "low NO_x " which probably mean different things to different readers.) As stated above, I believe that it may be the most important distinction among the various experimental results.
	Ans: ECHO 2003: $0.2 \le NO \le 1.3 \pmod{\text{data for:}} > 0.75 \operatorname{ppb}; 1.1 \le NO_2 \le 6.5 \pmod{\text{data for:}} > 2.5 \operatorname{ppb}$ ATTO 2015: $NO < 0.5 \operatorname{ppb}; NO_2 < 0.5 \operatorname{ppb}$ See also I_s as function of $NO \binom{NO_2, NO_x}{1}$ in Fig. S22.
49 - 49	Rev: I. 48-49: I disagree, a direct relation is shown in your equations (11) and (21). The isoprene flux is contained in two leading terms: one is in the TPI of the covariance budget and the other is in the variance budget (GPvar). But, as you say here, they may be more or less influential depending on the strength of other competing terms.
	Ans: A <u>direct</u> relation is not given because I_s is not directly proportional to the isoprene surface flux, as <u>other terms</u> of the corresponding equation significantly contribute to the result. (see section 4.1) In addition: The term in Eq. (9) respectively Eq. (11) (which may be related to the influence of the flux of isoprene on Is in Eq. (3) or Eq. (6)) is the <u>product</u> of the flux <u>and</u> the vertical derivative of the mixing ratio profile (Stull, 1988, p. 133). For an upward directed flux at the surface, in general the mixing ratio decreases with height, and, $\partial c_i/\partial z$ becomes negative. (Convention: An upward directed flux is a loss at the surface, and, therefore, has a negative sign). The (positive) influence of both product terms enter into the Eq. (11) for the variance of isoprene and even if all other terms would vanish, the influence of $\partial c_i/\partial z$ would remain.
	Therefore, the correlation between I_s and the flux decreases as given in Figures 10 and 11 already near the surface and even more in the ABL. Considering the balance of the variance, what is often called "a correlation with the flux of isoprene" is the correlation with the variance (or standard deviation) of isoprene.
51	Rev: I. 51: I realize that this may seem like quibbling, but chemical scalar fluxes do not necessarily always decrease with height. For example, if the entrainment is strong and the CBL concentration high enough, then it is possible for the isoprene vertical flux to increase with height (e.g. water vapor fluxes in some cases.)

	Ans: For isoprene the observations given by Su et al. (2016) for the time and area where Kaser et al. (2015) performed their measurements suggest decreasing fluxes with increasing height. For ECHO 2003 we have fluxes obtained near top of canopy and flux (profiles) decreasing with height (not published). Here we consider these data sets from measurements only. If entrainment of isoprene – as it is reported for water vapor – would be observed a secondary circulation must have transported isoprene above the inversion and chemical removal must be small, so that entrainment fluxes becomes significant (and advection would be the controlling term).
58	Ref: I. 58: You should probably be more quantitative about this statement. How much does Is increase with measurement bandwidth? It seems that if this is one of the leading findings of the study, worthy of inclusion in the abstract, then it should be elaborated a bit more: what does the cospectrum of <iso'oh'> look like at low wavenumbers? The increase in lower frequencies that you investigate in this study goes from about 2-10 km (based on a wind speed of about 4 m/s), comparable to the LES domain of Ouwersloot et al. (2011). On the other hand, the scales covered in Kaser et al. (2015) start at about 3 km (30 s OH measurements) and run to 50-100 km, which are dramatically larger scales than even your expanded analysis. This is the main reason that I believe the inclusion of the results of Kaser et al. (2015) is not appropriate for this work.</iso'oh'>
	Ans: We refer here (in the short abstract) to the data analysis given in Figs. 12 and 13, but possibly not explicitly enough explained in the text (line 1379 - 1386). We therefore, will add an additional description in a Supplement.
70	In general the contribution from large frequencies to any quantity increases with height. The spectral behavior of ogives (and the interaction of upward and downward transport) was discussed by Dlugi et al. (2014) for $cov(OH, ISO)$ in their Fig. 17. We performed the same analysis as for ECHO also for ATTO and will give examples on these results in a Supplement related to current Fig. 12 and Fig. 13 in the text. In the Abstract we will clarify the procedure; e.g.: "The spectral contribution to the covariance $cov(OH, ISO)$ and I_s was analyzed. Both quantities increase with increasing contribution of lower frequencies." Note that the frequency range by Kaser et al. (2015) is about $4.5 \cdot 10^{-4} Hz \le v \le 0.033 Hz$. It was artificially extended to higher frequencies by postulating a one to one proportionality between OH and O ₃ . We had $1.7 \cdot 10^{-3} Hz \le v \le 0.2 Hz$ (Dlugi et al. 2010 and 2014) with an extension to $5.6 \cdot 10^{-4} Hz \le v \le 0.2 Hz$ from the analysis of measured data as given in Table 2 / Fig. 12 -13 and the text below Fig. 12. Therefore, the difference is given mainly in the high frequency range, if no extension would be applied. This shows that a comparison of the data sets is not limited by significant differences in spectral contributions to $cov(OH, ISO)$ although the spatial scales covered are different. See also above comment.
70	Rev: I. 70: "and" is probably more often the case (buoyant & shear production combined). Ans: We will replace "and" by "as well as"
160	Rev: I. 160: missing 'g' in Tg. Ans: We correct this typing error (" T_a ")
165	Rev: I. 165: I don't see why the transport has to take place in a cloud-

	topped boundary layer. Ans: The experiments ECHO 2003 and ATTO 2015 were performed in a cloud topped ABL. In general, the notion "ABL" also includes clouds. Therefore we write: in the ABL
172	Rev: I. 172: I don't think that NO_2 generally determines the OH reactivity in any significant way. Also, whenever there is any significant isoprene, it tends to be the dominant VOC sink for OH. Therefore it doesn't make sense to consider isoprene moving through a static OH field because isoprene is usually the dominant sink and determines, in part, the OH field. Of course, this rapid reactivity is what drives the anticorrelation.
470	Ans: Here we give relevant reactants for the chemical interacting cycles not only directly related to OH reactivity. You write that isoprene determines, in part, the OH field. "In part", yes: Pfannerstill et al. (2018) published that the isoprene-related OH reactivity on the same day we made our measurements was 53 ± 29 % (average ± standard deviation) of the total OH reactivity. Therefore, the anti -correlation is not (-1) but significantly below (-0.5). OH is locally produced and destroyed with a chemical lifetime below 1 s, while isoprene has a lifetime of about 300 s or larger. Or in terms of Damköhler number OH has $Da > 20$ for a fast reacting compound while isoprene has $0.01 < Da < 0.5$. We do not consider a static OH field. But the locally variable OH is not transported on a scale above about several centimeters and the variability of OH is given by all chemical sources and sinks on the scales of this small volume. Therefore, we give this conceptual frame as suggested also in Dlugi et al. (2014).
178	Rev: I. 178: Doesn't Is < 0 hold only for when the chemical sink of OH is dominated by isoprene or something correlated with isoprene? If isoprene is correlated with a major source of OH (e.g. RO2 or HCHO) then Is could be > 0 in principle, no? Come to think of it, this is where I believe the answer to the origin of your observed Is limitation lies: the OH photochemistry is so heavily buffered that while isoprene is a dominant sink it is also correlated with important sources such as isoprene peroxy radicals, Iso-O2, and HCHO. See Kaser et al., 2015 Fig. S9 for estimates of the relative magnitudes of ROx (=RO2 + HO2) source strengths.
	Ans: With our measurements we try to find out "what can be seen in the atmosphere" or literarily spoken "what does nature tell us?". We thought about and discussed the segregation problem in general and in very detail for more than 20 years, but in the introduction of this paper we give the status of results being published. We consciously avoid in this context to mix information with qualitative world views. By definition $0 \le I_s \le 1$. But why $ -I_s < 0.3$ and not $ -I_s > 0.5$? Kaser et al. (2015) give box model results (well mixed conditions) on chemical sources and sinks for OH and suggest which pathways are important, as several (cited) researchers did before. They compared these "box results" to their measurements in a turbulent and convective atmosphere. They do not quantify the causes for turbulent fluctuations and the occurrence of $ I_s \gg 0$ but state that their data show high $ -I_s $ over areas with high computed surface fluxes. Here we refer to section 3.3 of Ouwersloot et al. (2011) where they showed that regions of higher isoprene emission may dynamically decouple from the surrounding (their case LSB2) resulting on high $I_s = -0.405$ but with other chemistry above the different spatial parts. Flying

	and weaker downdrafts resulting in a high $I_s \approx -0.3$ by averaging over different parts of the flight track (see Ouwersloot et al. (2011), page 10697). We did not discuss this topic because such conditions cannot be simply related to ECHO 2003/ ATTO 2015. But we will add some more results obtained by Ouwersloot et al. (2011) to our introduction and also to our graphical presentation if we present I_s as function of height.
188	Rev: I. 188: 'caused by' is an odd way to put it because Is < 0 is by definition represents an anticorrelation, but what actually caused it is the broader question that this paper tries to address.
	Ans: Yes, you are right. We changed the text accordingly.
207	Rev: I. 207: The Is values of Kim et al. (2016), albeit very small, nearly double across the range of NO _x from the experiments you compile (\sim 0.1 from ATTO to \sim 2 ppb from ECHO) from about -0.02 to -0.035.
	Ans: Regarding Fig. 2 of Kim et al. (2016) Is decreases with increasing NO_x (se also our fig. S22) and increases again for the very high NO_x case, where NO_x acts as an OH-sink. We included NO_x in our analysis (see Figs. S17 and S22 in the Supplement). From the measurements (ATTO and ECHO) there is no observable trend of Is with changing NO_x . We will add the intercomparison of near surface measurements related to ECHO 2003 and ATTO 2015 and also refer to Ouwersloot et al. (2011) and their findings in the revised manuscript.
209	Rev: I. 209: The UNO run of Ouwersloot et al. (2011) developed an Is of -0.12 and it was 'homogenous' in heat and isoprene fluxes, whereas without NO ₂ (lower NO _x) the control run Is = -0.07. Do you mean Is (low NO _x) > Is (high NO _x) or their magnitudes? Note that Ouwersloot et al. (2011) (from their Section 3.6) "stress the need to take the VOC and NO _x conditions into account in future studies that aim at segregation parameterizations." This advice seems to have been overlooked in the present work.
	Ans: The values and cases in this section have been given for comparison and to give some overview of the existing literature. Given the above discussion and the inclusion of the NO _x dependence in our revised manuscript we need to elaborate more on this point. Kim et al. (2016) found a slight decrease of Is in the surface layer by increasing the NO _x surface flux two times by one order of magnitude ("very low NO _x " to "High NO _x "). This decrease is attributed to higher mean OH that causes lower mean isoprene and therefore lower isoprene fluctuations. Another increase by a factor of 5 brings them to the "very high NO _x " case where NO _x becomes a significant OH sink and Is increases again. Ouwersloot et al (2011) state that a change of one order of magnitude in surface NO _x fluxes did not significantly change Is. The cases discussed here are with 0.5 ppb NO _x in the free troposphere, which is entrained into the BL and causes larger values of Is. These scenarios are different from the conditions given by Kim et al. (2016) and by our surface measurements. We include the discussion of the NO _x dependency into the revised manuscript. A first result is given in Fig. S22. The UNO run had $NO_2 = 0.5 ppb$ in the free troposphere and otherwise is comparable to run HOM. I _s (HOM) = -0.07 , I _s (UNO) = -0.124 . Adding spatial heterogeneity in isoprene source strength and moisture and heat fluxes (HNO) leads to I _s = -0.209 . Without any spatial variability of fluxes of moisture and heat and no NO_2 entrained from the free troposphere, but a spatial heterogeneity in the isoprene source

213	strength (LSB1) $I_s = -0.07$. Changing the scale of heterogeneity leads to $I_s(LSB2) = -0.405$, $I_s(LSB3) = -0.308$ and $I_s(LSB4) = -0.177$ for example. Here we refer to their discussion on the influence of enhanced NO_x in these model calculations in their section 3.6. Such conditions are completely different from our field studies ECHO 2003 / ATTO 2015 where NO_x sources are located mainly near or at the surfaces. In a Supplement (see Fig. S22) we add $ -I_s $ versus NO_x for ECHO and ATTO to present the results and to compare with literature and especially results from Kim et al. (2016).) Rev: I. 213: Again, I disagree with the statement that most of the results of Ouwersloot et al. (2011) are ≤ -0.1 From their Table 4
	HOM <is> = -0.07 which is definitely not < -0.1! (see less important weakness (1) above about comparisons of Is magnitude or numerical values less than zero.) Ans: Most results for homogeneous cases (and no gradients) are smaller or about $I_s = -0.1$ (<: "smaller than"). For clarification we may write instead $-I_s < 0.1$. Any "disturbance" to these cases cause $-I_s \ge 0.1$.</is>
243	Rev: I. 243: I think you should define this here: <w'c'2>. At this point I did not have any idea what M21 represented. Incidentally, this appears to be the best predictor you have observed to correlate with Is, so why not emphasize that more and show the results for M21 vs. Is in the ECHO & ATTO data sets? Ans: We will give the explicit notation of M₂₁ in the text and also add M₂₁ versus I_s graphical results for both studies near canopy top in another section of this paper. The corresponding figures are given in the</w'c'2>
	Supplement (Fig. S20, Fig. S21).
255 -257	Rev: I. 255-257: I think you should point out here that most theoretical treatments show Is to be of smaller magnitudes in the bulk of the CBL vs. the surface layer (e.g., Kaser et al., 2015; Ouwersloot et al., 2011; Patton et al., 2001.)
	Ans: In this section of the manuscript we only refer to Kaser et al. (2015): a) Patton et al. (2001) present results only near canopy top also as function of Da: $Da = 0.17$, $I_s = -0.05$; $Da = 0.6$, $I_s = -0.17$ with a very simplified chemistry.
	b) Kaser et al. (2015) give I_s as function of height throughout the ABL from model calculations (model used by Patton et al. (2001)) and found significantly larger values of I_s near canopy top, smaller ($0.08 \le -I_s \le 0.2$) near $z/z_i = 0.5$ in the ABL and larger also near z_i . c) Ouwersloot et al. (2011) give examples on $I_s(z)$ which show such results also
	In order to clarify this point we will add some comments on vertical profiles of $L(z)$ from model calculations in line 261
316	Rev: I. 316: The measurement of HOx fluxes and other higher moments
	from July 25 of the ECHO campaign seems like new material (not covered in Dlugi et al., 2014) so it probably merits some more explication. For instance, how did the cospectra compare to <w't'> or <w'o3'> or some other scalars? Does the flux direction and magnitude agree to theoretical predictions (e.g. Gao & Wesely, 1994)? Were these data analyzed in a separate paper? What changed to allow for these measurements, of which I know of no others? See Section 5.1 of Dlugi et al. (2014): "Spatial derivations of mixing ratios of these compounds and their fluxes are not available from this data set". What has changed?</w'o3'></w't'>

	And The floor of a disclosured as before a source of a disc Disc.
	et al. (2010) for information were not discussed here as this was not the scope of this manuscript which is already quite extensive. "What has changed"? Some of us in 2017 (M. Berger, M. Zelger, R. Dlugi, G. Kramm) (re-) analyzed additional calibrated data from ECHO 2003 to allow at least better estimates of terms in the budget equations, which were not evaluated in 2013 / 2014. We applied some of these results in this paper. Most results on reactive compounds and their fluxes from ECHO 2003 are not published, although they were evaluated. (see also line 740 - 746).
321 - 323	Rev: I. 321-323: What seems more directly important than temperature
	and humidity is the mean concentrations of HOx on that day relative to the rest of the experiment.
	Ans: Temperature and humidity conditions also influence chemistry. But we will add OH and NO_x - mixing ratios in the text and in the Supplement (see Fig. S10 – Fig. S13 and Fig. S15 – Fig. S16).
345	Rev: I. 345: If the HOx data is available at 0.2 Hz, why were there no
	fluxes reported other than July 25?
	Ans: As described in Dlugi et al. (2010), at the end of their section 1. "a
	one-day feasibility study was performed" and segregation could be
	calculated from these measurements for OH + isoprene, OH +
	monoterpenes the first time for atmospheric conditions and fluxes for
	HO_2 are given as well.
378	Rev: 1. 378: What was an average background OH value relative to the total? And as far as the second unit is concerned, how are you sure the amount of background OH is the same in both units? It seems that it might be worthwhile to describe some statistics of the backgrounds for both units to understand their variability and similarity.
	Figure 1: This figure is probably not necessary since none of the analysis of the OH measurements is covered herein.
	Ans: We give more details in the revised version. Fig. 1 will be completed to show distances between measuring volume of the eddy system and the different inlets which are needed for calculation of time shifts for determination of covariances and mixed higher moments.
453	Rev: I. 453: Equation 4 is a few pages later than this reference, but come to think of it equation 4 does not really give any chemical information other than there is a reaction between OH and isoprene, and as such is probably not necessary.
	Ans: Eq. (4) gives what is done in this paper, and, only therefore is given a number.
457	Rev: I. 457: The justification of time resolution invoked by Karl et al. (2013) is for turbulent fluxes (from Lenschow & Kristensen, 1985), which relies on assumptions about the turbulent statistics of w', but the requirements for a covariance with another scalar are different.
	Ans: This reference is given by Kaser et. al. (2015) in their paper and we cite only what is written in their paper and in their supplement. A discussion of their results in other sections is also only based on their reported findings.

489 -491	Rev: I. 489-491: I don't think it makes sense to mention compressible fluids and refer to the various 'a' variables defined by Richardson only to redefine them.
	Ans: We did it because to name the definitions clarifies the theoretical framework on which most atmospheric measurements are based. This is not trivial, but often forgotten. Segregation therefore is given in the "Reynolds" description of fluids, applied also in most atmospheric models. The averaging procedure and choice made to replace averages by time averages is given to refer to these aspects of "point" measurements which "see" some of the spatial variability of their surroundings. Higgens et al. (2013) only recently discussed this "Ergodic hypothesis" and compared it to experimental findings.
500	Rev: I. 500: I do not think that ''' is the best symbol to use for multiplication? It looks like a cross product and the subscripting i,j,k looks like tensor notation. I think a '×' or nothing at all is much more conventional and clear to indicate scalar multiplication.
	Ans: "x" is given by ACP for multiplication. The further notation follows the theoretic framework given – for example – by Vilá – Guerau de Arellano and Vinuesa or Verver in their modelling papers. ($i \equiv ISO$, $j \equiv OH$ and k, l, m, n are index for physical variables)
507	Rev: I. 507: Very rapidly after the H abstraction of the OH + isoprene reaction the production of some isomer of an Iso-O2 radical occurs. These peroxy radicals reacting with NO are usually a very important source of OH (Kaser et al., 2015 estimate it to be of similar magnitude as ozone photolysis.)
	Ans: We analyze the measured data in terms of Eq. (3) and further equations. Of this chemical cycle, the first step is analyzed and evaluated. If some compound is not measured, we cannot analyze it within this concept. We know that peroxy radicals exist, but only HO_2 was directly measured (and analysis of segregation - e.g. of $HO_2 + NO$ – has started)
508	Rev: I. 508: Is this supposed to mean Eq. 2 subtracted from Eq. 1? This was misleading it should be "and".
515	Rev: I. 515: What exactly is CxHyOz? I recognize the attempt to remain general, but I think the more important general species that is not mentioned anywhere is RO2.
	Ans: We give the general form, which includes RO2. : $RO_2 \cong C_x H_y O_2$ with $R = C_x H_y$
529	Rev: I. 529: the denominator of term 1 should be rwci.
574	Rev: I. 574: Using the Einstein summation convention with tensor notation confuses your own convention of using a generalized chemical trace gas, ci and cj. These look like 3D vectors in tensor notation (i, j = 1, 2, 3), and there is no other place in the manuscript where it is beneficial to generalize the chemistry. This is a work centered on the reaction of OH and isoprene and as such there is nothing gained by calling those species cj and ci, respectively. For example, kij, looks like a second order tensor, not a scalar reaction rate coefficient. Table 1: It would be more clear if you preserved the signs of these terms

	such that MR & TR are always <0 (that is, act to reduce <ci> in the budget.) Realizing that you labeled the terms inside the parentheses in equation (7), the terms DMF and DTF have different units because they are not the divergence thus the numbers in the table do not have to sum. I think it would be a lot clearer if you defined the terms to each entire item in the budget equation and keep their signs clear and indicative of how they change <ci>. Ans: Chemical compounds have index i, j, other (e.g., wind velocity vector components) have k, I, m following – for example - Vinuesa or Vilá – Guerau de Arellano. k_{ij} is explained in the text as all other quantities</ci></ci>
	Table 1: We will write – for example – (for S) - $-0.8 \cdot 10^{-3}$ to $1.2 \cdot 10^{-3}$, and all other quantities for clarity. In Eq. (7) we will write each term separated from each other, so that formally dimensions are the same. DMF and DTF are the spatial derivatives and dimensions are identical for all terms as for MR and TR
	and S
603 -628	Rev: I. 603-628: This entire paragraph suspiciously omits any mention of horizontal components of these advective terms. That, in itself, seems like an oversight, and it renders the last sentence (I.626-628) incorrect: a change in mean horizontal advection (without a change in the wind field divergence) can lead to significant changes in <ci>. It is not clear whether you mean 3D or 2D divergence/convergence in this discussion. Keep in mind that the second term in equation (8) in its full 3 dimensions is zero because of the incompressibility of the mean flow.</ci>
707	Ans: We add the discussion of horizontal advection influences. Rev: I. 707: Doesn't the concentration of isoprene decrease with height directly above the canopy making the numbers you report -0.01 to -0.07 ppb m-1? This term then is always negative acting to reduce the steady- state variance, no? This is counterintuitive, but is the nature of your steady-state approximation in equation (11).
	Ans: Yes, we introduce the sign for clarification. The term GP_{var} is given by the product of the flux and the corresponding "gradient". The source of heat, water vapour or isoprene is "at the surface". The surface (the plants at the surface) sees a loss as the flux is directed upwards. Therefore the convention is to consider this flux to have a negative sign (see – for example – the results given by R. Stull (1988 respectively 1993) in "An Introduction to Boundary Layer Meteorology" for variances of θ_v (page 113) or moisture (page 130). [For momentum the same concept holds.]
717	Rev: I. 717: I think you should be more specific in this reference. I believe it is specifically discussed in Section S3.2 of the Dlugi et al. 2014 supplementary materials.
	Ans: We will expand our discussion on this topic because Dlugi et al. (2010) described events with downdrafts <u>and</u> Dlugi et al. (2014) in more detail horizontal advection. We will give both situations with explanation of the specific conditions (and its possible influence on differences of mixing ratios ΔC) and the same estimates for ATTO 2015.
723	Rev: I. 733: Fig. 1 of Spirig et al. (2005) indicates a tower separation of
	Ans: Spirig et al. (2005) performed measurements at the west tower and

	the main tower (225 m away from the west tower). A third tower was installed in a distance of 120 m to the East from the main tower. So the correct number is 225 m – as determined by A. Schaub (co- author) to support the physical modelling by Aubrun (2005) instead of 125 m (typing error) in our text. The calculations were done with 225 m.
740 -746	Rev: I. 740-746: TTvar is the divergence of a turbulent flux of variance. Speaking of a "vertical change of TTvar" sounds like you are now looking at the second derivative of the variance flux. Is that what you're referring to? It would help if this discussion were a lot more clear about what is a turbulent flux of variance (the <w'c'2> term), vs. its vertical change (TTz,var).</w'c'2>
	Ans: This was indeed misleading and we refer here to the difference in the variance flux with height. We got a proportionality between both terms for isoprene and Θ , and, on two days also results for two heights above canopy for ISO, Θ , and q. These results are applied to estimate TT_{var} for isoprene. We will give a more detailed explanation in the revised version (Supplement). We clarified the text. (see also answer to comment on line 316).
752	Rev: I. 752: Your term III in Eq (11) is equivalent to IV3 in Table 4 of Dlugi et al. (2014) which states its estimated magnitude as < 3e-5 ppb2 s-1.
	Ans: We did a reanalysis of all these data for ECHO 2003 and corresponding calculations for ATTO 2015. But indeed the given exponent is wrong for both experiments: $< 10^{-5} ppb^2 s^{-1}$ is the correct information as not all values are below $10^{-6} ppb^2 s^{-1}$.
762	Rev: I. 762: It is not clear how you estimate the gradient of a fluctuating scalar directly, but in general variance budget discussions the molecular destruction term is expected to be first order (to balance mean gradient production in the steady-state, flow-integrated condition.) See Section 5.3 of Wyngaard (2010), for example.
	Ans: The c_i (fluctuations) were measured on the given days for several hours in two heights above canopy with time resolution of 1 Hz and synchronization of better than 0.005 s. The equation yields correct dimensions: $10^{-5}m^2s^{-1} \cdot (ppb^2m^{-2}) = 10^{-5}ppb^2s^{-1}$ (see also answer to comment on line 316).
805	Rev: I. 805: How did you derive these OH flux values? And are you imply that you have these values for both experiments? Is this not discussed anywhere else in the literature? It seems like a very difficult measurement to directly make by eddy covariance. In any event, you should probably specify the sign of this flux (I believe it should be downward, <0). These magnitudes seem much larger than predicted by Gao & Wesely (1994).
	Ans: In section 2 we describe that we measured time resolved OH mixing ratios (0.2 Hz for ECHO respectively 0.067 Hz for ATTO) and wind velocity components u, v, w and could filter the w- series to 0.2 Hz
	respectively 0.067 Hz for calculate $w'OH'$ for both field studies. OH- fluxes are explained in this section and they cannot be applied to calculate a deposition flux, they are given only by term S and MR(TR) in Eq. (7) by the influences of chemical sinks and sources on S and MR(TR). The effective distance for physical transport is just a few cm but the

	numerical values are caused by the source and sink distributions of OH. (In line with Gao and Wesly 2004).
834 -835	Res: I. 834/5: Again, the gradients of isoprene should be negative.
	Ans: Yes, we give always the sign in the revised version.
845	Res: I. 845: I recommend sticking to a single format for all of these range limits of your scale analysis, and preferably using only one significant digit. For example, 'xe-3' to 'ye-1'. Two significant digits for these scale analyses that typically span multiple decades just seems unnecessary and slightly confusing.
	Ans: The referee is right and we will give only one digit in the revised manuscript.
859 -861	Res: I. 859-861: The similarity you are applying to associate the different scalar transport terms needs to be explicitly stated. It seems like you are using some sort of modified Bowen ratio analog to the transport term, but this seems highly speculative. I believe that speculative is fine, but it would be more convincing if you explicitly stated the similarity you are applying.
	Ans: Our description needs to be clarified: We will express especially our own data and give the argument clearer in terms that the heat flux and the turbulent transport of the heat flux are calculated from measurements and are compared to results from measurements of isoprene in two heights above canopy on some days (see above) to calculate comparable terms. A proportionality is obtained for the term III for heat and ISO which is used for this estimate. Term IV is calculated as described in the text. See also answer to comment on line 316).
893	Rev: I. 893: This range of a factor of 5 for the pressure transport term implies that the time scale values have a range of a factor of 6, because the isoprene fluxes mentioned above span a factor of 30 (0.02 to 0.6 ppb m/s). It would be clearer if you presented what the mixing length concept of Poggi et al. (2004) depended on.
	Ans: We will add a short description of this mixing length concept (see Poggi, 2004) in the revised version.
973	I. 973: I have tried and tried and redone the arithmetic on the governing equation (15), because I know how pernicious and elusive sign errors can be, but I just cannot see how the normalized variance term in equations (20 & 21) can have the opposite sign of the Cij term (which is the balance of the terms from Rij outside of the covariance and variance terms that all have the same sign). This same error appears in Dlugi et al. (2014) at their equation (15). This has very important bearing on the analysis because the normalized variance of isoprene and the RES (Eq. 16) terms both act to increase the magnitude of the OH and isoprene segregation coefficient, in this case, - Is, because Is < 0. It seems like this equation will change the authors' calculations of REis because they solve for it as the residual of equation (21), and will fundamentally change Figure 7.
	Ans: We thank the reviewer very much for identifying this sign error by his careful analysis. Starting with Eq. (15) a sign was incorrectly transferred so that Eq. (21) has a wrong sign. The correct form reads $I_s + CH_{is} + nvar(ISO)_{is} + RE_{is} = 0$. As mentioned in the text for the revised figures the presentations are

	made with $I_s + CH_{is} + nvar(ISO)_{is} = -RE_{is}$. The revised Fig. 5 – Fig. 7 also correct results presented by Dlugi et al. (2014) with respect to the magnitude of RE_{is} and the relation between I_s and $nvar(ISO)_{is} = (-RE_{is})$.
999	Rev: I. 999: You say that Rij goes negative despite terms (b) and (c) which are positive definite. But Rij is defined with a negative (definite) sign (equation 18), so the positive definite terms like (b) and (c) work to make Rij negative. I found this language error typical throughout the manuscript. When revising I recommend being very careful with the language about discussing relative values or magnitudes of values, always retaining the accurate signs of the terms.
	Figure 3: 60% of graph has no information on it. Also, why is the total term in one unit (ppb2 s-1) and the individual components in another (ppb3). I think it makes the figure less clear to include the reaction rate in one and eliminate it in the others. Figure 4: Again, why compare these terms of differing units and then put a one-to-one line on the figure? Also why ignore the sign of Rij? If all the values are negative, then label it –Rij
	Ans: R_{ij} can be positive or negative and may change sign (Fig. 3) for ATTO 2015. The revised Fig. 3 is given in the Appendix. The text has been clarified according to the comment.
1041	Rev: I. 1041: Term (c) is not the only leading term of Rij. The ATTO results could differ because of a substantially different contribution from the <oh'iso'>[Iso] term (a), especially at higher values of [OH]var(Iso).</oh'iso'>
	Ans: The revised Fig. 3 shows that term c is dominant together with term a. (The latter is used to formulate a diagnostic Eq. (21) for I_s). As term a) is negative and term c) is positive (for ATTO) their difference determines the presentation in Fig. 3 with partly different results for ATTO and ECHO.
1109	Rev: I. 1109: "Rij increases [in magnitude] with increasing variance"
1126	Rev: I. 1126: It does not seem clear from Fig. 9 that the relationship between rij and Is is non-linear. You use this term a lot but none of the figures clearly show any distinction among a linear or non-linear relationship.
1041	Rev: I. 1041: Term (c) is not the only leading term of Rij. The ATTO results could differ because of a substantially different contribution from the <oh'iso'>[Iso] term (a), especially at higher values of [OH]var(Iso).</oh'iso'>
	Ans: The revised Fig. 3 shows that term c is dominant together with term a. (The latter is used to formulate a diagnostic Eq. (21) for I_s). As term a) is negative and term c) is positive (for ATTO) their difference determines the presentation in Fig. 3 with partly different results for ATTO and ECHO.
1109	Rev: I. 1109: "Rij increases [in magnitude] with increasing variance"
1106	Ans: We will add "in magnitude"
1126	between rij and Is is non-linear. You use this term a lot but none of the

	figures clearly show any distinction among a linear or non-linear relationship.
	Ans: A numerical fit to the data will be added to show this relation.
1219 - 1221	Rev: I. 1219-1221: You do not know for certain that the var(Iso) and flux terms are only established near the surface (for example, the entrainment zone can possess high variances and fluxes.) Furthermore, equation (11) also shows that var(Iso) is augmented by a term proportional to [Isoprene] <ci×cj> (concordant with equation (21) with the corrected sign), which could also be a leading term near the surface. Also note that what you are referring to as GPvar actually serves to decrease isoprene variance in the steady-state form you present in equation (11) because d[Iso]/dz < 0. In the variance budget, equation (9), GPvar produces variance, but in the reactive chemical steady-state of equation (11) it reduces variance.</ci×cj>
	Ans: We discuss the related terms with respect to the application to our field studies near the surface and some discussion of results from literature (e.g. Kaser et al. 2015). The upward directed "fluxes" have a negative sign (they describe a sink at the surface; see definition in Stull (1988)). Also $(\partial c_i/\partial z) < 0$ (as the reviewer states), so that the product is positive in both Equations. We refer to the height dependence of different terms as cited in lines 1205 – 1212.
1243	Res: I. 1243: I am assuming you mean vertical advection by the mean flow. However, just because W is larger in magnitude at higher elevation in the CBL does not mean that the magnitude of the scalar gradient is larger. It is much more likely to be turbulent transport that is a large term. If by 'vertical advection' you mean turbulent transport (the divergence of a vertical turbulent flux), then I would specify that.
	Ans: Here we refer to arguments given to your comments line $1331 - 1333$ (see below). In addition the term A_{var} in Eq. (9) (respectively Eq. (11)) can be of influence on the magnitude of variance.
1244	Rev: I. 1244: Is is related to the isoprene flux by two separate terms of Eq. (21): the TPI term of REis and the GPvar term in the normalized variance, nvar(Iso). This is not made clear in this discussion and consequently these arguments are ambiguous. These two flux terms have different coefficients (OH and isoprene gradients, respectively) so that their coefficients will change with altitude (probably both decreasing with height.) I would suggest eliminating all of this height dependence of variance discussion because it is speculative (for reactive scalars) and it does not really help the overall work in any way that I can discern.
	Ans: Here we refer to arguments given to your comments line $1331 - 1333$ (see below). In addition: we applied the convention that upward directed fluxes are a sink of isoprene (or heat, or moisture) at the surface and a gain for the atmosphere. Therefore, upward directed fluxes have a (-) sign. Thus, the term GP_{var} becomes positive in Eq. (9) as well as in Eq. (11).
1319	Rev: I. 1319: No, OH and O3 do not necessarily have a large positive covariance (presumably someone could check if there were O3 fluxes being measured on the tower), but the principal source of OH (on the ~1 s time scale) is the photo dissociation of O3 so it is very likely that they are, in fact, correlated.

	Ans: Yes, we agree; here we only repeat the arguments originally given by Kaser et al. (2015). To clarify we quote the text in "".
1331 - 1333	Rev: I. 1331-1333: That is patently incorrect. First, Ei0 is directly related to the flux at any height in the CBL (you used such a relationship yourself earlier to extrapolate their observed fluxes at $z/zi \sim 0.4$ to the surface). Furthermore, as stated previously, Is is correlated to the isoprene flux through both the GPvar (where it serves to diminish the variance, and thus Is , right above the canopy where the flux is upward and the gradient is negative), and in the TPI term of REis in (21) where it tends to be a source of negative covariance because the OH gradient is likely positive (due to preponderance of sinks effusing out of the canopy.)
	Ans: As described in the text the relation termed "no longer valid" is that Is can be approximated by the isoprene (surface) flux, not that the isoprene flux at a certain level is not related to the surface flux (this was just one condition to be meet to establish this relationship). The whole paragraph describes that one can establish such relationships, but that the predictive power of the surface values diminishes with increasing height in the ABL as other terms become more important in the budgets. We applied the CBL scaling as described in lines 1195 – 1203. This concept considers the ABL a slab with boundaries at bottom and at top and the emission flux at bottom and the entrainment flux at top. An emission flux (like the sensible heat flux) decreases with height – represented like a flux divergence; the sensible heat flux crosses the zero line, if entrainment is observed (see some summary in Sorbjan (1989)). Here the change of mixing ratio with time is proportional to the flux
	divergence as given in the balance of the mixing ratios (Eq. (7)). Thus the flux decreases with height. In line $1205 - 1221$ we focus on term GP_{var} from the balance of variance, because of Eq. (3) – the definition of I_s . We argue that the contribution of the term, which is the product of the isoprene flux and the (vertical) gradient, decreases with height, and therefore the correlation coefficient obtained at the surface between the isoprene flux and I_s decreases with increasing height. Therefore, findings given by Kaser et al. (2015) and in our paper are consistent. The term TPI in RE_{is} in Dlugi et al. (2014) is the product of the isoprene flux and the spatial gradient of the OH mixing ratio. At least the isoprene flux decreases with height. Above canopy top we observed a decrease of OH with height for ECHO 2003. For ATTO 2015 we have no data on the vertical OH profile. At least for ECHO this influence on I_s decreases with height. Due to your remarks we will relate the discussion of GP_{var} and TPI in the revised version.
	Rev: Figure 12: Why are there are not the same number of blue diamonds (spectrally extended) as there are black circles? They should be 18-27% larger according to line 1391. Also, the blue diamonds all lie exactly on top of the circles showing no spectral change in <oh'iso'>. Also, the two points on the lower left (Is < -0.2) do not seem to exist on Fig. 14.</oh'iso'>
	Ans: The calculation of I_s for 30 minutes intervals instead of 10 minutes intervals is related to the calculation of the covariance but also the calculation of the product of the means. Both quantities changed compared to the results for 10 minute intervals resulting in less values of

	the covariance. If results for 30 minutes overlap with a value obtained for 10 minutes this does not mean that the related values (belonging to the 30 minutes interval, $3x10 = 30$) are not larger by the given percentage increase. The overlapping was caused by the choice of resolution of the graphic software, which was set too low (see revised fig. 12). We corrected this so that "overlapping" is avoided and explain this in the text below that figure. The blue dashed curve gives the empirical range for all data.
1409 – 1410	Rev: I. 1409-1410: Is never becomes independent of [Iso][OH] because that product resides in its denominator. The covariance may become independent, but not Is.
	Ans: For covariances smaller than $-4x10^{-5}$ (larger negative numerical value) $mean(OH) \cdot mean(ISO)$ stays approximately constant. Therefore the increase of I_s becomes only controlled by the covariance. We clarified the text accordingly.
1460	Rev: I. 1460: According to Dlugi et al. (2014) Eq. (17) M12 were considered the "ejections", and M21 the "sweeps"?
	Ans: This was a typing error. Furthermore, we add the results of Is versus M_21 as Fig. S20 – Fig S21 to the Supplement.
	Rev: Figure 14: This figure is nearly identically the same as Dlugi et al. (2014) Figure 20, save for the three modeling results and two Kaser et al. (2015) points. Why do you not present any of the ATTO data on this figure? Why plot both BP and kinematic heat flux? As far as I can discern there is no appreciable difference in the underlying relationship and plotting both just clutters the figure.
	Ans: We add the ATTO data to this presentation (see Fig. 14b - bottom). The value $BP > 3 \cdot 10^{-3} m^2 s^{-3}$ is the range of the onset of free convective conditions. We added this discussion to the text. This is a physical criterion, so no statistical comparison is made because we sort data (of I_s) in terms of the surface sensible heat flux and the buoyant production term.
1545	Rev: I. 1545: It is very challenging to find an empirical relationship in Fig. 14 as stated. You should propose one if you think it exists. Is looks to me like a nearly vertical line would fit through the points of BP > 3e-3? I wonder what the p-value of such a fit would be, because it does not look great by my eye.
	Ans: There is no direct correlation of Is versus heat flux expected, but as thermals are known to have a profound influence on Is it is important to know if data points were derived in a convective regime or not. This graph is not meant to find a predictor, but to sort the data (see also general comments). If we consider the concept of the shear and buoyant generation (or production) we noticed that a certain empirical relation exists between the buoyant production BP and I_s (and the sensible heat flux and I_s). For small BP, I_s is not dependent on BP, but if a value of $3 \cdot 10^{-3} m^2 s^{-3}$ is approached I_s – values increase. Here the data are in the range where the free convection limit is approached, which is about $3 \cdot 10^{-3} m^2 s^{-3}$ following for example results given by Stull (2000). Model results from

	results from ATTO 2015 increase if this "limit" is approached, but become smaller than $ -I_s = 0.04$ for $BP > 3 \cdot 10^{-3} m^2 s^{-3}$. In this range the correlation coefficient in Eq. (3) becomes small (see also the revised figures Fig. 14a, Fig. 14b and Fig. 9).
1597 -1601	 Rev: I. 1597-1601: If M21 vs. nvar(Iso) & REis shows a strong relationship as in Fig. 18 of Dlugi et al. (2014) why not show it? If this finding is worthy of a paragraph in conclusion, then it seems it should be represented in a figure. Earlier you state the sweeps only weakly correlate with nvar(Iso) and REis, and here you state that only ejections contribute to Is. This all seems to beg for a figure of both M21 and M12 to see how much they each correlate to nvar(Iso) & REis. This could be a micrometeorological parameter that is readily measured in canopy field studies that could be used to estimate Is for chemical modelers, for example. Ans: We give I_s versus M₂₁ in the Supplement figures (Fig. S20 and Fig. S21)
1621	Rev: I. 1621: The bandwidth of the Kaser et al. (2015) measurements were out to nearly 100 km. For typical winds speeds of, say, 5 m/s this would require a 5.5 hr integration time at a tower site. Thus the measurements, aside from being made several hundred meters higher than the ECHO & ATTO datasets, represent a much larger spectral band. The 'hypothesis' of scale dependence is established explicitly in Ouwersloot et al. (2011), why bring this in as a hypothesis from this work? There is currently no easy way to disentangle the isoprene surface source variability from the scale of the measurements in terms of their effects on Is, so it is not a hypothesis that is truly tested in this work.
	Ans (see also general comments): Kaser et al. (2015) give horizontal length scales from about 3 km up to 220 km. With the mean measuring velocity of the aircraft of 100 m s ⁻¹ one obtains their frequency range of measured data (OH) between $4.5 \cdot 10^{-4}$ Hz and 0.033 Hz. Note that we obtained data for $1.7 \cdot 10^{-3}$ Hz (respectively $5.4 \cdot 10^{-4}$ Hz) and 0.2 Hz (ECHO) respectively 0.067 Hz (ATTO). Kaser et al. (2015) extrapolated their spectra for OH by a direct proportionality to the spectra of 03 (not to the chemical production term of OH, which would include the fluctuations of photolysis frequency and water vapor). By this procedure they extended their spectral presentation to higher frequencies. On average they give $Is \sim 0.13$ with a contribution of the range >5 $\cdot 10^{-3}$ Hz by absolute values of -6% to -8%
	The scales in space are not of interest, only the ranges of frequencies f or wave numbers $k=2\pi f u'$ (u= Wind velocity or flight velocity) are important. Therefore the data given by Kaser et al. (2015) directly fit into our Fig. 12. The results from Ouwersloot et al. (2011) are taken from their original LES data from mean vertical profiles for the lowest layers to compare to our results from heights above rough surfaces. The Ergodic theorem tells us that one can compare averages over space with averages over time.
1625 -1627	Rev: I. 1625-1627: This is an interesting idea, but not very well explicated in the body text of the manuscript, and only sprung on the reader in the last sentence of the work. The diurnal source correlations (which in and of itself would promote a positive species covariance) occur on long time scales relative to the chemistry and the TKE dissipation and the 10-40 minutes averaging used in this study. In order for this to be a reason for the "limits" of Is suggested on the 10 min scale the sources would need to correlate on this restricted time scale, and/or

there would need to be some sort of downscale cascade at play. This speaks to the absence of any cospectral representation of Is in this work (something like Fig. S4 of Kaser et al., 2015), which would help understand its spectral dependence. In any event, I suspect the compensating chemistry of OH sources that are correlated with isoprene (e.g. isoprene peroxy radicals) are the most likely culprits for limiting the magnitude of Is.
Ans: This part of the "summary and conclusions" will be revised. The "restricted time scale" is only one aspect, but if extended, the absolute value of I_s increases. If we consider the results obtained by Ouwersloot et al. (2011), the strategy how to sample over up- and downdrafts becomes important. We analyze "ogives" to present these results in the text (with information also in the supplement), to give limits to I_s from the mixing processes. In this context we will also discuss possible limits to I_s occurring from chemical cycling.