



Abstract

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27 We present a comparison of 1064 nm aerosol optical depth (AOD) and aerosol extinction

28 profiles from the Cloud-Aerosol Transport System (CATS) Level 2 aerosol product with

29 collocated Aerosol Robotic Network (AERONET) AOD, Aqua and Terra Moderate Imaging

30 Spectroradiometer (MODIS) Dark Target (AOD) and Cloud-Aerosol Lidar with Orthogonal

31 Polarization (CALIOP) AOD and extinction data for the period of Feb. 2015-Oct. 2017. Upon

32 quality assurance checks of CATS data, reasonable agreements are found between aerosol data

33 from CATS and other sensors. Using quality assured CATS aerosol data, for the first time,

34 variations in AODs and aerosol extinction profiles are evaluated at 00, 06, 12, and 18 UTC

35 (and/or 0:00 am, 6:00 am, 12:00 pm and 6:00 pm local solar times) on both regional and global

36 scales. This study suggests that marginal variations are found in AOD from a global mean

37 perspective, with the maximum and minimum aerosol vertical profiles found at local noon and

38 6:00 pm local time respectively, for both the June-November and December-May seasons.

39 Strong diurnal variations are found over North Africa and India for the December-May season,

40 and over North Africa, Middle East, and India for the June-November season. In particular, over

41 North Africa, during the June-November season, a diurnal peak in aerosol extinction profile of

42 20% larger than daily mean is found at 6:00 am (early morning local time), which may possibly

43 be associated with dust generation through the breaking down of low level jet during morning

44 hours.

45

46



1.0 Introduction

Aerosol measurement through the sun-synchronous orbits of Terra and Aqua by nature encourages a larger scale, daily average point of view. Yet, we know that pollution (e.g., Zhao et al., 2009; Tiwari et al., 2013; Kaku et al., 2018), fires and smoke properties (e.g., Reid et al., 1999; Giglio et al., 2003; Hyer et al., 2013), and dust (e.g., Mbourou, et al., 1997; Fielder et al., 2013; Heinold et al., 2013) can exhibit strong diurnal behavior. Sun-synchronous passive satellite aerosol observations from the solar spectrum only provide a small sampling of the full diurnal cycle and geostationary sensors such as the Advanced Himawari Imager (AHI) on Himawari 8 (Yoshida et al., 2018) and Advanced baseline Imager on GOES-16/17 (Aerosol Product Application Team of the AWG Aerosols/Air Quality/Atmospheric Chemistry Team, 2012) satellites, while an improvement over their predecessors, must overcome the broader range of scattering and zenith angles (Wang et al., 2003; Christopher and Zhang, 2002) with no nighttime retrievals. AEROSOL ROBOTIC NETWORK (AERONET; Holben et al., 1998) based sun photometer studies improve sampling, but until very recently with the development of a prototype lunar photometry mode, are also limited to daylight hours. The critical early morning and evening are largely missed in solar observation based approaches.

Observation-based diurnal variations of aerosol properties are needed for improving chemical transport modeling, geochemical cycles and ultimately climate. The measurement of diurnal variations of aerosol properties resolved in the vertical is especially crucial of aerosol phenomena for visibility and particulate matter forecasts. Indeed, the periods around sunrise and sunset show significant near surface variability that is difficult to detect with passive sensors. While lidar data from Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) provide



69 early afternoon and morning observations, two temporal points and a 16 day repeat cycle are
70 insufficient to evaluate the morning and evening hours.

71 Some of the limiting factors in previous studies can be addressed by the Cloud-Aerosol
72 Transport System (CATS) lidar flying aboard the International Space Station (ISS) since 2015
73 (McGill et al. 2015). The ISS's precessing orbit with a 51.6° inclination allows for 24 hour
74 sampling of the tropics to mid-latitudes, with the ability to observe aerosol and cloud vertical
75 distributions at both day and night time with high temporal resolution. For a given location
76 within $\pm 51.6^\circ$ (Latitude), after aggregating roughly 60 days of data, near full diurnal cycle of
77 aerosol and cloud properties can be obtained from CATS observations (Yorks et al. 2016). This
78 provides a new opportunity for studying diurnal variations (day and night) in aerosol vertical
79 distributions from space observations.

80 Use of CATS has its own challenges. Most importantly, CATS retrievals must cope with
81 variable solar noise around the terminator where we expect the strongest diurnal variability to
82 exist. Further, CATS lost its 532 nm channel early in its deployment, leaving only a 1064 nm
83 channel functioning. The availability of only one wavelength limited the CATS cloud-aerosol
84 discrimination algorithm, which can cause a loss of accuracy compared to CALIPSO which has
85 2 wavelengths. This deficiency is in part overcome by using the Feature Type Score. Using two
86 years of observations from CATS, in this paper, we focus on understanding of the following
87 questions: How well do CATS derived aerosol optical depth (AOD) and aerosol vertical
88 distributions compare with aerosol properties derived from other ground-based and satellite
89 observations such as AERONET, MODIS and CALIOP? Do differences exhibit a diurnal cycle?
90 What are the diurnal variations of aerosol optical depth on a global domain? What are the
91 diurnal variations of aerosol vertical distribution on both regional and global scales?



92

93 **2.0 Datasets**

94 Four datasets, including ground-based AERONET data, as well as satellite retrieved
95 aerosol properties from MODIS and CALIOP, are used for inter-comparing with AOD and
96 aerosol vertical distributions from CATS. Upon thorough evaluation and quality assurance
97 procedures, CATS data are further used for studying diurnal variations of AOD and aerosol
98 vertical distributions for the period of Feb. 2015 – Oct. 2017.

99

100 **2.1 CATS**

101 CATS Level 2 (L2) Version 2-01 5 km Aerosol Profile products (L20_D-M7.2-V2-
102 01_5kmPro, L20_N-M7.2-V2-01_5kmPro) were used in this study for the entire period of CATS
103 operation on the ISS (~Feb. 2015–Oct. 2017). CATS L2 profile data are provided at 5 km along-
104 track horizontal resolution and 533 vertical levels at 60 m vertical resolution and a wavelength of
105 1064 nm. CATS also provides data at 532 nm, but due to a laser-stabilization issue, 532 nm data
106 is not recommended for use (Yorks et al. 2016). Thus, only 1064 nm products were used in this
107 study. CATS data are quality-assured following a manner similar to Campbell et al. (2012),
108 which was applied to CALIOP. QA thresholds (including extinction QC flag, Feature Type
109 Score, and uncertainty in extinction coefficient) are listed below:

- 110 (a) Extinction_QC_Flag_1064_Fore_FOV is equal to 0
- 111 (b) Feature_Type_Fore_FOV = 3 (aerosol only)
- 112 (c) $-10 \leq \text{Feature_Type_Score_FOV} \leq -2$
- 113 (d) $\text{Extinction_Coefficient_Uncertainty_1064_Fore_FOV} \leq 10 \text{ km}^{-1}$



114 Extinction was also constrained to the nominal range provided in the CATS data catalog
115 ($0 \leq \text{Extinction_Coefficient_1064_Fore_FOV} \leq 1.25 \text{ km}^{-1}$), similar to several previous
116 studies (Redemann et al., 2012; Toth et al., 2016). Due to the presence of a large amount of near
117 zero negative extinction values, profiles containing these values were also considered by setting
118 all near zero negative values to zero, resulting in an effective threshold of
119 $\text{Extinction_Coefficient_1064_Fore_FOV} \leq 1.25 \text{ km}^{-1}$. For this study, both the
120 $\text{Aerosol_Optical_Depth_1064_Fore_FOV}$ and $\text{Extinction_Coefficient_1064_Fore_FOV}$ datasets
121 were used to provide AOD and 1064 nm extinction profiles (hereafter the term “extinction” will
122 refer to 1064 nm unless explicitly stated otherwise), respectively.

123

124 2.2 CALIOP

125 NASA’s CALIOP is an elastic backscatter lidar that operates at both 532 nm and 1064
126 nm wavelengths (Winker et al., 2009). Being a part of the A-Train constellation (Stephens et al.,
127 2002), CALIOP provides both day- and night-time observations of Earth’s atmospheric system,
128 at a sun-synchronous orbit, with a narrow swath of around 70 m and a temporal resolution of ~16
129 days (Winker et al., 2009). For this study, CALIOP Level 2.0 5 km Aerosol Profile products
130 (L2_05kmAProf) are used for inter-comparing to CATS retrieved AODs and aerosol vertical
131 distributions.

132 L2_05kmAProf data are available at 5 km horizontal resolution along-track and include
133 aerosol retrievals at both 532 nm and 1064 nm wavelengths. The vertical resolution is 60 m
134 near-surface, increasing to 180 m above 20.2 km in MSL altitude. As only 1064 nm CATS data
135 are used in this study as mentioned above, likewise only those CALIOP parameters relating to
136 1064 nm are used in this study (Vaughan et al., 2018; Omar et al., 2013). Note that as suggested



137 by Rajapakshe et al. (2017), lower single to noise ratio (SNR) and higher minimum detectable
138 backscatter are found for the CALIOP 1064 nm data in-comparing with the CALIOP 532 nm
139 data. Also, the CALIOP aerosol layers are detected at 532 nm and the 1064 nm extinction is
140 only computed for the bins within these layers. This may introduce a bias for aerosol above
141 cloud studies. In this study, Extinction_Coefficient_1064 and
142 Column_Optical_Depth_Tropospheric_Aerosols_1064 are used for CALIOP extinction and
143 AOD retrievals, respectively (Vaughan et al., 2018; Omar et al., 2013). As with the CATS data,
144 CALIOP data are quality-assured following the quality assurance steps as mentioned in a few
145 previous studies (e.g. Campbell et al., 2012; Toth et al., 2016; 2018). These QA thresholds are
146 listed below:

- 147 (a) Extinction_QC_Flag_1064 is equal to 0,1,2,16, or 18
- 148 (b) Atmospheric_Volume_Description = 3 (aerosol only)
- 149 (c) $-100 \leq \text{CAD_Score} \leq -20$
- 150 (d) $\text{Extinction_Coefficient_Uncertainty_1064} \leq 10 \text{ km}^{-1}$

151 Furthermore, as in Campbell et al. (2012), only those profiles with AOD > 0 were
152 retained in order to avoid profiles composed of only retrieval fill values. Extinction was also
153 constrained to the nominal range provided in the CALIOP data catalog ($\text{Extinction_1064} \leq 1.25$
154 km^{-1}), similar to our QA procedure for CATS as described above.

155

156 **2.3 MODIS Collection 6.1 Dark Target product**

157 Moderate Resolution Imaging Spectroradiometer (MODIS) Aqua and Terra Collection
158 6.1 Dark Target over-ocean AOD data (Levy et al., 2013) were used for comparison to CATS
159 AOD. The data field of “Effective_Optical_Depth_Best_Ocean” were used and only those data



160 flagged as “good” or “very good” by the Quality_Assurance_Ocean runtime QA flags are
161 selected for this study, similar to Toth et al. (2018). Because MODIS does not provide AOD in
162 the 1064 nm wavelength, AOD retrievals from 860 and 1240 nm spectral channels are used to
163 interpolate AODs at 1064 nm, assuming a constant angstrom exponent as suggested by Shi et al.,
164 (2011; 2013). Only totally cloud free (or cloud fraction equal to zero) retrievals, as indicated by
165 the Cloud_Fraction_Land_Ocean parameter are used.

166

167 **2.4 AERONET**

168 By measuring direct and diffuse solar energy, AERONET observations are used for
169 retrieving AOD and other ancillary aerosol properties such as size distributions (Holben et al.,
170 1998). AERONET data are considered as the ground truth for evaluating CATS retrievals in this
171 study. Only cloud screened and quality assured version 3 level 2 AERONET data at the 1020
172 nm spectrum are selected and are used for inter-comparing with CATS AOD retrievals at the
173 1064 nm wavelength. AERONET does not have specific guidance on error in the 1020 nm
174 channel, as it is known to have some thermal sensitivities. However they do report significantly
175 more confidence in version 3 of the data, which has temperature correction (Giles et al., 2018).
176 Error models are ongoing, and for this study we assume double the RMSE, or +/-0.03.

177

178 **3.0 Results & Discussion**

179 **3.1 Inter-comparison of CATS data with AERONET, MODIS and CALIOP data**

180 In this sub-section, the performance of over land and over ocean CATS AOD retrievals
181 are compared against AERONET and C6.1 over ocean MODIS DT aerosol products. CATS



182 derived aerosol extinction vertical distributions are also cross-compared against collocated
183 CALIOP aerosol extinction vertical distributions.

184

185 3.1.1 CATS-AERONET

186 As the initial check, CATS data from Feb. 2015-Oct. 2017 are spatially (within 0.4
187 degree Latitude and Longitude) and temporally (± 30 minutes) collocated against ground-based
188 AERONET data. Note that one AERONET measurement may be associated with several CATS
189 retrievals in both space and time, and vice versa. Thus, both CATS and AERONET data are
190 further averaged spatially and temporally, which results in only one pair of collocated and
191 averaged CATS and AERONET data for a given collocated incident. Also, only data pairs with
192 AOD larger than 0 from both instruments are used for the analysis. This step is necessary to
193 exclude CATS profiles with all retrieval fill values as discussed in Section 2 (Toth et al., 2018).
194 Note that the CATS-AERONET comparisons are for daytime only, and higher uncertainties are
195 expected for CATS daytime than night AODs.

196 As shown in Figure 1a, without quality-assurance procedures, high spikes in CATS AOD
197 of above 1 (1064 nm) can be found for collocated AERONET data with AOD less than 0.3 (1020
198 nm). Those high spikes in CATS AOD may due to cloud contamination in the V2-01 CATS
199 daytime data, which will be improved in the upcoming CATS V3-00 data products. Upon
200 completion of the QA steps as outlined in Section 2.1, a reasonable agreement is found between
201 quality-assured CATS (1064 nm) vs. AERONET (1020 nm) AODs with a correlation of 0.64
202 (Figure 1b). Comparing Figure 1a with 1b, with the loss of only ~10% of collocated pairs due to
203 the QA procedures, we have observed an overall improvement in correlation between CATS and
204 AERONET AOD from 0.17 to 0.64. Note that similar results are found in comparisons between



205 collocated CATS and MODIS/CALIOP data without the use of QA procedures on CATS data.
206 Thus, only QAed CATS data are used hereafter. Still, this exercise highlights the need for
207 careful quality checks of the CATS data before applying the CATS data for advanced
208 applications to overcome cloud-aerosol discrimination uncertainties.

209

210 3.1.2 CATS-MODIS

211 To examine over ocean performance, column integrated CATS AODs are inter-compared
212 with collocated Terra and Aqua C6.1 MODIS DT over ocean AOD, interpolated to 1064 nm.
213 Over ocean C6.1 MODIS DT data are selected due to the fact that higher accuracies are reported
214 for over ocean versus over land MODIS DT AOD retrievals (Levy et al., 2013). In addition,
215 comparing with over land MODIS DT data, which provides AOD retrievals at three discrete
216 wavelengths (0.46, 0.55 and 0.65 μm), over water AOD retrievals are available from 7
217 wavelengths including the 0.87 and 1.24 μm spectral channels, allowing a comparison with
218 CATS AOD at the same wavelength upon interpolation.

219 MODIS and CATS AOT retrievals are collocated for the study period of Feb. 2015-Oct.
220 2017 (Figure 2). Pairs of CATS and MODIS data were first selected for both retrievals that fall
221 within ± 30 minutes and 0.4 degrees latitude and longitude of each other. Then, similar to the
222 AERONET and CATS collocation procedures, collocated pairs were further averaged to
223 construct one pair of collocated MODIS and CATS data for a given collocation incident. Shown
224 in Figure 2a, a correlation of 0.71 is found between collocated over water Terra MODIS C6.1
225 DT and CATS AODs with a slope of 0.78. Similar results are found for the comparisons
226 between over water Aqua MODIS and CATS AODs with a correlation of 0.75 and a slope of
227 0.79.



228

229 **3.1.3 CATS-CALIOP AOD**

230 In the previous two sections, AODs from CATS are inter-compared with retrievals from
231 passive-based sensors such as MODIS and AERONET. In this section, AOD data from
232 CALIOP, which is an active-based sensor, are evaluated against AOD retrievals from CATS.
233 Again, for each collocation incident, pairs of CALIOP and CATS data are selected in which both
234 retrievals fall within ± 30 minutes temporally and 0.4 degrees latitude and longitude spatially.
235 There could be multiple CATS retrievals corresponding to one CALIOP data point, and vice
236 versa. Thus, the collocated pairs are further averaged in such a way that only one pair of
237 collocated CATS and CALIOP data is derived for each collocation incident.

238 Figure 3a shows the comparison of CATS and CALIOP AODs for all collocated pairs
239 including both day- and night-time. A reasonable correlation of 0.7, with a slope of 0.69, is
240 found for a total of 2681 collocated data pairs. Further breaking down the comparison into day
241 and night cases, a much better agreement is found between the two datasets during nighttime
242 with a correlation of 0.84 and 0.81 for over-ocean and over-land cases respectively. In
243 comparison, a lower correlation of 0.62 (0.52), with a slope of 0.44 (0.63), is found between the
244 two datasets, using over land (ocean) daytime data only, for a total of 171 (1207) collocated
245 pairs. This result is not surprising as daytime data from both CALIOP and CATS are expected to
246 be noisier due to solar contamination (e.g. Omar et al., 2013; Toth et al., 2016).

247 Still, larger discrepancies between CATS and CALIOP AODs during daytime indicate
248 that both sensors are more susceptible to solar contamination. To overcome solar contamination
249 and more accurately detect aerosol layers, CALIOP and CATS data products are averaged up to
250 80 km and 60 km, respectively. Noel et al. (2018) found that clouds screened using the feature



251 type score were accurately detected by CATS data products throughout the diurnal envelope of
252 solar angles. To ensure the solar contamination does not introduce a diurnal bias in aerosol
253 detection or products, CATS AODs are further evaluated as a function of local time. For each
254 CATS observation of a given location and UTC time, the associated local time is computed by
255 adding (subtracting) the UTC time by 1 hour per 15° Longitude away from the Prime Meridian
256 in the east (west) direction. Figure 4a shows the CATS AOD versus local time for both global
257 land and oceans. While noisy in data, an averaged AOD peak is found around local noon that is
258 about 0.02-0.03 higher than both sunrise and sunset times. Still, for high AOD cases, no
259 significant solar noon peak is found. Also, no major deviations in AODs are found during either
260 sunrise or sunset time, although it is speculated that larger uncertainties in CATS AODs and
261 extinctions may be presented around day and night terminators. Figure 4b shows a similar plot
262 as Figure 4a, but with the region restricted to 25°S-52°S. Here, we want to investigate the
263 variations in CATS AODs as a function of local time, over relatively aerosol free oceans. We
264 picked 25°S as the cutoff line as CATS data only available to 51.6°S (limited to the ISS
265 inclination angle) and thus, this threshold is used to ensure enough data samples in the analysis,
266 although some land regions are also included. As indicated in Figure 4b, a clear diurnal variation
267 is found, with the peak mean AOD of 0.08 found around local noon and smaller AOD values of
268 0.06 found for both sunrise and sunset times. Also, no significant deviations in pattern are found
269 for both sunrise and sunset time, plausibly indicating that solar contamination, as speculated,
270 may not be as significant. It is, however, unclear if the 0.02 AOD difference between local noon
271 and sunrise and sunset times is introduced by retrieval bias or indeed a physical existence.

272 To further explore the 0.02 difference, Figure 4c shows the difference between
273 AERONET (1020 nm) and CATS (1064 nm) AOD (Δ AOD) as a function of local time, again,



274 although data are rather noisy, no major pattern is found near sunrise or sunset times, again,
275 further indicating that solar contamination during dawn or dusk times, may have a less severe
276 impact to CATS AOD retrievals from a long term mean perspective.

277 In summary, Sections 3.1.1-3.1.3 suggest that with careful QA procedures, AOD
278 retrievals from CATS are comparable to those from other existing sensors such as AERONET,
279 MODIS, and CALIOP at the same local times.

280

281 **3.1.4 CATS-CALIOP Vertical Extinction Profiles**

282 One advantage of CATS is its ability to retrieve both column-integrated AOD and
283 vertical distributions of aerosol extinction. Therefore, in this section, extinction profiles from
284 CATS are compared with that from CALIOP. Again, similar to the Section 3.1.3, collocated
285 profiles for CATS and CALIOP are first found for both retrievals that are close in space and time
286 (within ± 30 minutes and 0.4 degrees latitude and longitude). However, different from Section
287 3.1.3, only one pair of collocated CATS and CALIOP profiles, which has the closest Euclidian
288 distance on the earth's surface, is retained for each collocated incident.

289 The CATS cloud-aerosol discrimination (CAD) algorithm is a multidimensional
290 probability density function (PDF) technique that is based on the CALIPSO algorithm (Liu et al.
291 2009). The PDFs were developed based on Cloud Physics Lidar (CPL) measurements obtained
292 during over 11 field campaigns and 10 years. Figure 5 shows that CATS V2-01 aerosol
293 extinction agrees very well with CALIOP for nighttime (Figure 5c) and over land (Figure 5e).
294 However, CATS overestimates aerosol extinction around 1 km compared to CALIOP during
295 daytime (Figure 5b) and over ocean (Figure 5d). Based on statistical comparisons of CATS L2O
296 V2-01 cloud and aerosol detection frequencies with CALIOP, it was determined that, during



297 daytime over ocean, depolarizing liquid water clouds in the lower troposphere are sometimes
298 classified as lofted dust mixture or smoke aerosols in the CATS V2-01 data products. This is
299 primarily a result of enhanced depolarization ratios within liquid water clouds due to multiple
300 scattering (which is not represented in the CPL measurements used for the PDFs). To overcome
301 this issue, the CATS V3-00 CAD algorithm uses horizontal persistence tests and additional tests
302 using variables such as the perpendicular ATB, to better differentiate clouds and aerosols. More
303 details will be provided in an upcoming paper (Yorks et al., in prep). Since the CATS V3-00 data
304 has not been released yet, we will focus our discussion of aerosol diurnal variability on regions
305 primarily over land.

306 CATS also has a stronger extinction when compared to CALIOP in the lowest 2 km,
307 which may be due to differences in cloud screening. Vertical profiles of collocated CATS and
308 CALIOP extinction for daytime only profiles and nighttime only profiles are shown in Figure 5b
309 and 5c, respectively. Compared to a total collocated pair count of 2681 in the overall profile
310 data, day and night profiles have 1342 and 1339 collocated pairs, respectively. Again, the shapes
311 of the CATS and the CALIOP nm extinction vertical profile are very similar for all three cases,
312 despite the above mentioned offsets in altitude. Figure 5d and 5e show the mean of those
313 extinction profiles which occurred over-water and over-land, as defined by the CATS surface
314 type flag. Again in both cases CATS and CALIOP have very similar shapes in their vertical
315 extinction profiles. The vertical structure of over-water extinction is also very similar to that of
316 all profiles, day, and night, which is perhaps not surprising as water profiles made up 2111 of
317 2681 (~79%) collocated pairs. The vertical structure of over-land is more different than the other
318 groups, as the extinction is higher throughout a larger depth of the atmosphere, tapering off much



319 more slowly from the surface. Furthermore, the extinction from CATS is actually lower than
320 CALIOP for over-land profiles, unlike all other categories.

321

322 **3.2 Diurnal Cycle of AODs and Aerosol Vertical Distributions**

323 Using the QAed CATS data, seasonal variations as well as diurnal variations in CATS
324 AODs are derived in this section. Diurnal variations in the vertical distributions of CATS aerosol
325 extinction are also examined at both global and regional scales.

326

327 **3.2.1 Seasonal and Diurnal Variation of AOD**

328 Figures 6a-b show the spatial distributions of CATS AODs at the 1064 nm spectral
329 channel for boreal winter-spring (Dec.-May, DJFMAM) and boreal summer-fall (June-Nov,
330 JJASON) seasons, for the period of Feb. 2015-Oct. 2017. To construct Figures 6a and 6b,
331 quality-assured CATS AODs are first binned on a 5 degree by 5 degree grid over the globe for
332 the above mentioned two bi-seasons. For each $5\times 5^\circ$ (Latitude/Longitude) bin, for a given
333 season, CATS AODs are averaged on a pass-basis first, and then further averaged seasonally to
334 represent AOD value of the given bin.

335 In DJFMAM season, significant aerosol features are found over North Africa, Mid-East,
336 India and Eastern China. For the JJASON season, besides the above mentioned regions, aerosol
337 plumes are also observable over Southern Africa, related to summer biomass burning of the
338 region (e.g. Eck et al., 2013). The seasonal-based spatial distributions of AODs from CATS,
339 although reported at the 1064 nm channel which is different from the 550 nm channel that is
340 conventionally used, are similar to some published results (e.g. Lynch et al., 2016).



341 For comparison purposes, Figures 6c-6d shows similar plots as Figures 6a-6b, but with
 342 the use of CALIOP AOD at the 1064 nm spectral channel. Note that those are climatological
 343 means rather than pairwise comparisons. While patterns are similar in general, at regions with
 344 peak AODs of 0.4 or above for CALIOP, such as North Africa for the DJFMAM season and
 345 North Africa, Middle-East and India for the JJASON, much lower AODs are found for CATS.
 346 In some other regions, such as over South Africa and upper-portion of Middle-East for the
 347 JJASON season, however, higher CATS AOD values are observed. Figures 6e and 6f show the
 348 similar spatial plots as Figures 6a and 6b but with the use of Aqua MODIS AODs from the DT
 349 products. For the Aqua MODIS DT products, aerosol retrievals at the short-wave Infra-red
 350 channels are only available over oceans, and thus Figures 6e-6f show only over ocean retrievals.
 351 Again, while general AOD patterns look similar, discrepancies are also visible, such as over the
 352 coast of south east Africa for the JJASON season. Those discrepancies may result from biases
 353 in each product, but it is also possibly due to the differences in satellite overpass times, as both
 354 MODIS and CALIOP provide early morning and afternoon over passes while CATS is able to
 355 report atmospheric aerosol distributions at multiple times during a day. It is also possibly due to
 356 aerosol above cloud related issues as reported by Rajapakshe et al. (2017).

357 Similar to Figures 6a and 6b, Figures 7a and 7b show the spatial distribution of CATS
 358 AODs, but for CATS extinction values that are below 1 km AGL only, for the DJFMAM and
 359 JJASON seasons respectively. Figure 7c and 7d (7e and 7f) show the CATS mean AOD plots
 360 for extinction values from 1-2 km AGL (> 2 km AGL). For the DJFMAM season, elevated
 361 aerosol plumes with altitude above 2 km AGL are found over the North coast of Africa. For the
 362 JJASON season, elevated dust plumes (> 2 km AGL) are found over North Africa and the
 363 Middle-East regions, while elevated smoke plumes are found over the west coast of South Africa



where above cloud smoke plumes are often observed during the Northern hemispheric summer season (e.g. Alfaro-Contreras et al., 2016).

CATS has a non-sun-synchronized orbit, which enables measurements at near all solar angles. Thus, we also constructed $5\times 5^\circ$ (Latitude/Longitude) gridded seasonal averages (for DJFMAM and JJASON seasons) of CATS AODs at 0, 6, 12 and 18 UTC that represent 4 distinct times in a full diurnal cycle, as shown in Figure 8. To construct the seasonal averages, observations within ± 3 hours of a given UTC time as mentioned above are averaged to represent AODs for the given UTC time. On a global average, the mean AODs are 0.090, 0.090, 0.090 and 0.091 for 0, 6, 12 and 18 UTC respectively for the JJASON season and are 0.101, 0.100, 0.097 and 0.097 for the DJFMAM season. Thus, no significant diurnal variations are found on a global scale, as global means are dominated by background aerosols that have weak diurnal variations in measured absolute AOD values.

Still, strong diurnal variations with the maximum averaged diurnal AOD changes of above 0.15 can be observed for regions with significant aerosol events such as Northern Africa and India for the DJFMAM season and Northern Africa, Southern Africa, Mid-East and India for the JJASON season, as illustrated in Figure 9. Note that Fig. 9a (9b) shows the maximum minus minimum seasonal mean AODs for the four difference times as shown in Figs. 8a,c,e,g (8b,d,f,h). Interestingly but not unexpectedly, regions with maximum diurnal variations match well with locations of heavy aerosol plumes as shown in Figures 6 and 8.

3.2.2 Diurnal variations of Aerosol Extinction on a Global Scale (both at UTC and local time)



386 Using quality-assured CATS derived aerosol vertical distributions, mean global CATS
387 extinction vertical profiles are also generated as shown in Figure 10. Similar to steps as
388 described in the section 3.2.1, CATS extinction profiles are binned into 00, 06, 12, and 18 UTC
389 times based on the closest match in time for the JJASON and DJFMAM seasons. Figure 10a
390 (10d) shows the daily averaged CATS extinction profiles in a black line, and 00, 06, 12 and 18
391 UTC averaged in blue, green, yellow and red lines respectively, for the DJFMAM (JJASON)
392 season. CATS extinction profiles for the daily average as well averages for the four selected
393 times are similar, suggesting that minor temporal variations in CATS extinctions can be expected
394 for global averages.

395 Those global averages are dominated by CATS profiles from global oceans (Figure 10b
396 and 10e), which also have small diurnal variations, as ~70% of the globe is covered by water. In
397 comparison, noticeable diurnal changes in aerosol vertical distributions are found over land as
398 shown in Figure 10c and 10f. For the DJFMAM season, at the 1 km altitude, the minimum and
399 maximum aerosol extinctions are at 12 and 18 UTC respectively. Similarly, the minimum and
400 maximum aerosol extinctions are at 18 and 6 UTC at the altitude of 400 m. For the JJASON
401 season, the minimum aerosol extinction values are found at 12 UTC for the whole 0-2 km
402 column, while the maximum aerosol extinction values are at 18UTC for 1.5 km and 0UTC for
403 the 300-400 m altitude. Still, it should be noted that aerosol concentrations may be a function of
404 local time, yet for a given UTC time, local times will vary by region. Also, due to solar
405 contamination, nighttime retrievals from CATS are considered to be less noisy than daytime
406 retrievals, and this difference in sensor sensitivity between day and night may further affect the
407 derived diurnal variations in CATS AOD and aerosol vertical profiles as shown in Figure 3 for
408 individual retrievals. Still, no apparent solar pattern is detectable from Figure 8, and only minor



409 diurnal variations are found for Figure 10a and 10d, which indicate that such a solar
410 contamination may introduce noise but not bias to daytime aerosol retrievals, from a global mean
411 perspective.

412 If we examine the mean global CATS extinction vertical profiles with respect to local
413 time as shown in Figure 11, however, some distinct features appear. For example, Figure 11a
414 and 11d suggests that on global average, the minimum and maximum aerosol extinction below 1
415 km is found for 6:00 pm and 12:00 pm local time, respectively for both JJASON and DJFMAM
416 seasons. Similar patterns are also observed for over global oceans. However, for over land
417 cases, for both seasons, peak in aerosol extinction is found at the 500-1000 m layer for local
418 noon, which is ~20-30% higher than daily mean values. This may indicate stronger solar heating
419 at the surface and hence stronger near surface convection at local noon that brings near surface
420 aerosol particles to a higher altitude.

421

422 **3.2.3 Diurnal variations of Aerosol Extinction on a Regional Scale (at local time)**

423 In this section, the diurnal variations of aerosol vertical distributions are studied as a
424 function of local solar time for selected regions with high mean AODs as highlighted in Figure 6.
425 We picked local solar time here as for those regional analyses, near 1 to 1 transformation can be
426 achieved between UTC and local solar time. Also, as learned from the previous section, aerosol
427 features are likely to have a local time dependency. A total of four regions, including Africa-
428 north, Middle East, India and Northeast China, which show significant season all mean AODs in
429 Figure 6, are selected for the DJFMAM season (Figure 12). For the JJASON season (Figure 13),
430 in addition to the above mentioned 4 regions, the Africa-south region is also included due to
431 biomass burning in the region during the Northern Hemisphere summer time. The



432 Latitude/Longitude boundary of each selected region is described in Table 1. Regional-based
433 analyses are also conducted for 4 (5) selected regions for the DJFMAM (JJASON) season at four
434 local times: 0:00 am, 6:00 am, 12:00 pm and 6:00 pm, using quality assured CATS profiles.
435 Generally, the maximum diurnal change in aerosol extinction is found at the altitude of below 1
436 km for all regions as well for both seasons. Also, larger diurnal variations in vertical
437 distributions of aerosol extinction are found for the JJASON season, in-comparing with the
438 DJFMAM season, while regional-based differences are apparent.

439 For the Africa-north region, dominant aerosol types are dust and smoke aerosol for the
440 DJFMAM season, and is dust for the JJASON season (e.g. Remer et al., 2008). Interestingly, the
441 maximum aerosol extinction below 500m is found at 6:00 am for the DJFMAM season. While
442 for the JJASON season, the maximum aerosol extinctions are found at 6:00 am for the whole 0-2
443 km column, with a significant ~20% higher aerosol extinction from either daily mean or vertical
444 profiles from 0:00 am, 12:00 pm and 6:00 pm. Note that 6:00 am in the Africa-north region
445 corresponds to early morning, which has been identified in several studies (Fiedler et al., 2013;
446 Ryder et al. 2015) as the time of day when nocturnal low-level jet breakdown causes large
447 amounts of dust emission in this region. Thus, we suspect that this large 6:00 am peak in
448 maximum aerosol extinctions may be the signal resulting from the low-level jet ejection
449 mechanism captured on a regional scale. As the day progresses into the afternoon and early
450 evening, we find the aerosol heights shifting upwards, likely related to the boundary layer's
451 mixed layer development.

452 For the Middle East region, for the JJASON season, a daily maximum in aerosol
453 extinction of $\sim 0.13 \text{ km}^{-1}$ is found at local morning or early morning (0:00 am and 6:00 am), with
454 a daily minimum of $\sim 0.09 \text{ km}^{-1}$ found at local noon (12:00 pm), for the peak aerosol extinction



layer that has a daily mean aerosol extinction of $\sim 0.11 \text{ km}^{-1}$. This translates to a $\sim \pm 20\%$ daily variation for aerosol extinction for the peak aerosol extinction layer. Much smaller daily variation in aerosol extinction, however, is found for the same region for the DJFMAM season.

For the India region, for the JJASON season, a large peak in aerosol extinction of up to 20% higher than daily mean is found at 6:00 am below 1 km. The minimum aerosol extinction is found at 0:00 am for the layer of $\sim 400\text{--}1000 \text{ m}$, and is overall $\sim 10\%$ lower than the daily means. The minimum aerosol extinction is found at 6:00 pm for the layer below 400 m. For the DJFMAM season, minimum aerosol extinctions are found at 12:00 pm for near the whole 0–2 km column, while for the layer below 500 m, the maximum aerosol extinction values are found at early morning (0:00 am and 6:00 am). This is consistent with the diurnal formation of significant haze.

For the Northeast China region, less diurnal variation is found for the DJFMAM season. Yet, a significant peak found at 1 km for local noon (12:00 pm) for the JJASON season, which is $\sim 30\%$ higher than daily averages for the JJASON season. The reason for this elevated peak at regional local noon, however, is not known, although it may relate to the peak in surface Particulate Matter concentrations. Lastly, for the Africa-south region, biomass burning aerosols are prevalent during the summer time and thus only the JJASON season is analyzed. As shown in 13b, below 500 m in altitude, lower extinction values are found for local afternoon (12:00 pm and 6:00 pm) and higher extinction values are found for local morning or early morning (0:00 and 6:00 am). Still, the diurnal variation in aerosol vertical distribution is rather marginally for the region.

4.0 Conclusions



478 Using CALIOP, MODIS and AERONET data, we evaluated CATS derived AODs as
479 well as vertical distributions of aerosol extinctions for the study period of for Feb. 2015 – Oct.
480 2017. CATS data (at 1064 nm) are further used to study variations in AODs and aerosol vertical
481 distributions diurnally. We found:

482 (1) Quality assurance steps are critical for applying CATS data in aerosol related
483 applications. With a 10% data loss due to QA steps, an improvement in correlation
484 from 0.17 to 0.64 is found for the collocated CATS and AERONET AOD
485 comparisons. Using quality assured CATS data, reasonable agreements are found
486 between CATS derived AODs and AODs from CALIOP, Aqua MODIS DT and
487 Terra MODIS DT at the same local times, with correlations of 0.70, 0.75 and 0.71
488 respectively.

489 (2) While the averaged vertical distributions from CATS compare reasonably well with
490 that from CALIOP, differences in peak extinction altitudes are present. This may due
491 to contamination of daytime aerosol detections over ocean by marine boundary layer
492 clouds in the CATS V2-01 data products, which will hopefully be resolved in the
493 future CATS V3-00 data.

494 (3) From the global mean perspective, minor changes are found for AODs at four
495 selected times, namely 00, 06, 12 and 18 UTC. Yet noticeable diurnal variations in
496 AODs of above 0.15 (at 1064 nm) are found for regions with extensive aerosol
497 events, such as over North Africa, and India for the DJFMAM season, and over North
498 and South of Africa, India and Middle East for the JJASON season.

499 (4) From the global mean perspective, changes are less noticeable for the averaged
500 aerosol extinction profiles at 00, 06, 12 and 18 UTC. Yet, if the study is repeated



501 with respect to local time, a peak in aerosol extinction is found for local noon and the
502 minimum value in aerosol extinction is found at 6:00 pm local time for both JJASON
503 and DJFMAM seasons. In particular, for over land cases, in both seasons, a lifted
504 aerosol plume at 500-1000 m altitude (with the peak aerosol extinction that is ~20-
505 30% higher than daily averages) is found at local noon, which may indicate the impact
506 of strong surface solar heating as well as stronger near surface convection on aerosol
507 vertical distributions.

508 (5) Larger diurnal variations are found at regions with heavy aerosol plumes such as
509 North and South (summer season only) of Africa, Middle East, India and Eastern
510 China. In particular, aerosol extinctions from 6:00 am over North Africa are ~20%
511 higher than daily means as well other three times for the 0-2 km column for the
512 JJASON season. We suspect this may be related to increase in dust concentrations
513 due to breakdown of low level jets at early morning time for the region.

514 This paper suggests that strong regional diurnal variations exist for both AOD and
515 aerosol extinction profiles. These results demonstrate the need for global aerosol measurements
516 throughout the entire diurnal cycle to improve visibility and particulate matter forecasts as well
517 as studies focused on aerosol climate applications.

518

519 **Author Contribution:**

520 Authors J. Zhang, J. S. Reid and L. Lee designed the study. L. lee worked on data processing for
521 the project. J. E. Yorks guided L. lee on data processing. The manuscript was written with
522 inputs from all coauthors.



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529



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Table 1. Geographic ranges, height above ground level of maximum extinction, diurnal extinction range at height of maximum extinction, and time (local) of peak extinction for the boxed red regions in Figure 6 and vertical profiles shown in Figures 12 and 13.

DJFMAM/JJASON					
Region	Latitude	Longitude	Height AGL (m) of Max. Extinction	Extinction Range (km^{-1}) at Height AGL of Max. Extinction	Time of Peak Extinction at Height AGL of Max. Extinction
India	7.5°N - 32.5°N	65°E - 85°E	180/240	0.109-0.131/0.138-0.182	6 am/6 am
Africa - North	2.5°N - 22.5°N	35°W - 20°E	420/480	0.107-0.130/0.098-0.121	12 pm/6 am
Africa - South	17.5°S - 2.5°N	0° - 30°E	/420	/0.090-0.100	/6 am
Middle East	12.5°N - 27.5°N	35°E - 50°E	240/180	0.093-0.116/0.081-0.135	6 am/0 am
China	27.5°N - 37.5°N	110°E - 120°E	240/240	0.107-0.154/0.085-0.133	6 am/6 am



Figure Captions

Figure 1. Collocated AERONET 1020 nm AOT vs. CATS 1064 nm AOD a) without CATS QA applied, and b) with CATS QA applied.

Figure 2. Collocated MODIS C6.1 a) Terra and b) Aqua estimated 1064 nm AOD vs. CATS 1064 nm AOD with CATS QA applied.

Figure 3. Collocated CALIOP 1064 nm AOD vs. CATS 1064 nm AOD with CATS QA applied for a) both day and night, b) nighttime over-land, c) nighttime over-water, d) daytime over-land, e) daytime over-water.

Figure 4: CATS 1064 nm AOD a) as a function of local time for the globe, and b) as a function of local time for areas south of -25 degrees. The difference between CATS 1064 nm AOD and AERONET 1020 nm AOD as a function of local time is shown in c). The mean is represented by the blue line, while the median is the green line.

Figure 5. CATS and CALIOP vertical profiles of 1064 nm extinction for a) all profiles, b) daytime only, c) nighttime only, d) over-water, and e) over land.

Figure 6. Mean AOD (1064 nm) by season for a) DJFMAM CATS, b) JJASON CATS, c) DJFMAM CALIOP, d) JJASON CALIOP, e) DJFMAM MODIS Aqua, and f) JJASON MODIS Aqua. Red boxes indicate locations of regional vertical distributions in Figures 12 and 13.

Figure 7. Mean CATS AOD (1064 nm) by season for a) DJFMAM below 1 km AGL, b) JJASON below 1 km AGL, c) DJFMAM 1-2 km AGL, d) JJASON 1-2 km AGL, e) DJFMAM above 2 km AGL, and f) JJASON above 2 km AGL.

Figure 8. Seasonal Mean AOD (1064 nm) binned by every 6-hours for a) DJFMAM 0 UTC, b) JJASON 0 UTC, c) DJFMAM 6 UTC, d) JJASON 6 UTC, e) DJFMAM 12 UTC, f) JJASON 12 UTC, g) DJFMAM 18 UTC, and h) JJASON 18 UTC.

Figure 9. Maximum minus minimum mean seasonal AOD (1064 nm) for a) DJFMAM, and b) JJASON.

Figure 10. Global mean 6-hourly vertical profiles of CATS 1064 nm extinction for a) DJFMAM all profiles, b) DJFMAM water profiles, c) DJFMAM not-water profiles, e) JJASON all profiles, f) JJASON water profiles, g) JJASON not-water profiles.

Figure 11. Global mean 6-hourly local time (0:00 am, 6:00 am, 12:00 pm and 6:00 pm) vertical profiles of CATS 1064 nm extinction for a) DJFMAM all profiles, b) DJFMAM water profiles, c) DJFMAM not-water profiles, d) JJASON all profiles, e) JJASON water profiles, f) JJASON not-water profiles.

Figure 12. DJFMAM 6-hourly average (local time; 0:00 am, 6:00 am, 12:00 pm and 6:00 pm) vertical profiles of CATS 1064 nm for locations shown in Figure 6a; a) Africa-north, b) Middle East, c) India, and d) Northeast China.



710 **Figure 13.** JJASON 6-hourly average (local time; 0:00 am, 6:00 am, 12:00 pm and 6:00 pm)
711 vertical profiles of CATS 1064 nm for locations shown in Figure 6b; a) Africa-north, b) Africa-
712 south, c) Middle East, d) India, and e) Northeast China.

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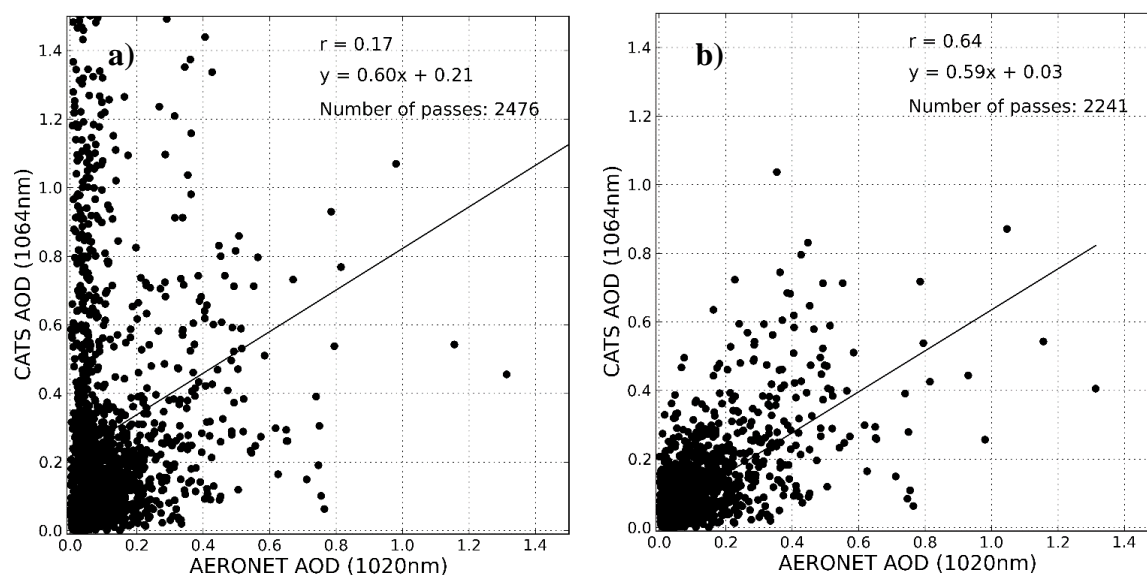


Figure 1. Collocated AERONET 1020 nm AOT vs. CATS 1064 nm AOD a) without CATS QA applied, and b) with CATS QA applied.



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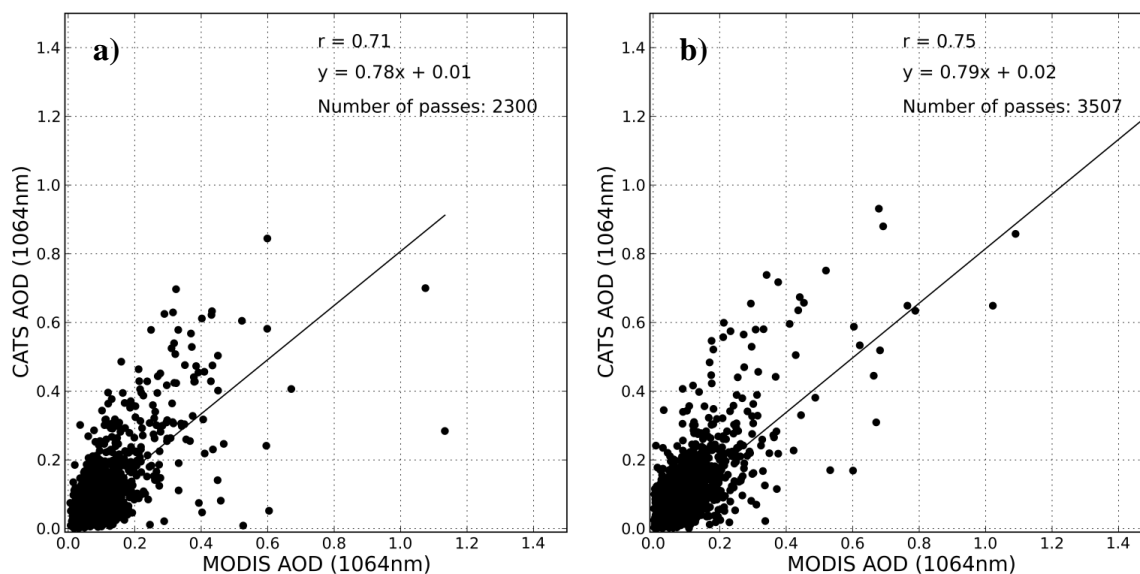


Figure 2. Collocated MODIS C6.1 a) Terra and b) Aqua estimated 1064 nm AOD vs. CATS 1064 nm AOD with CATS QA applied.



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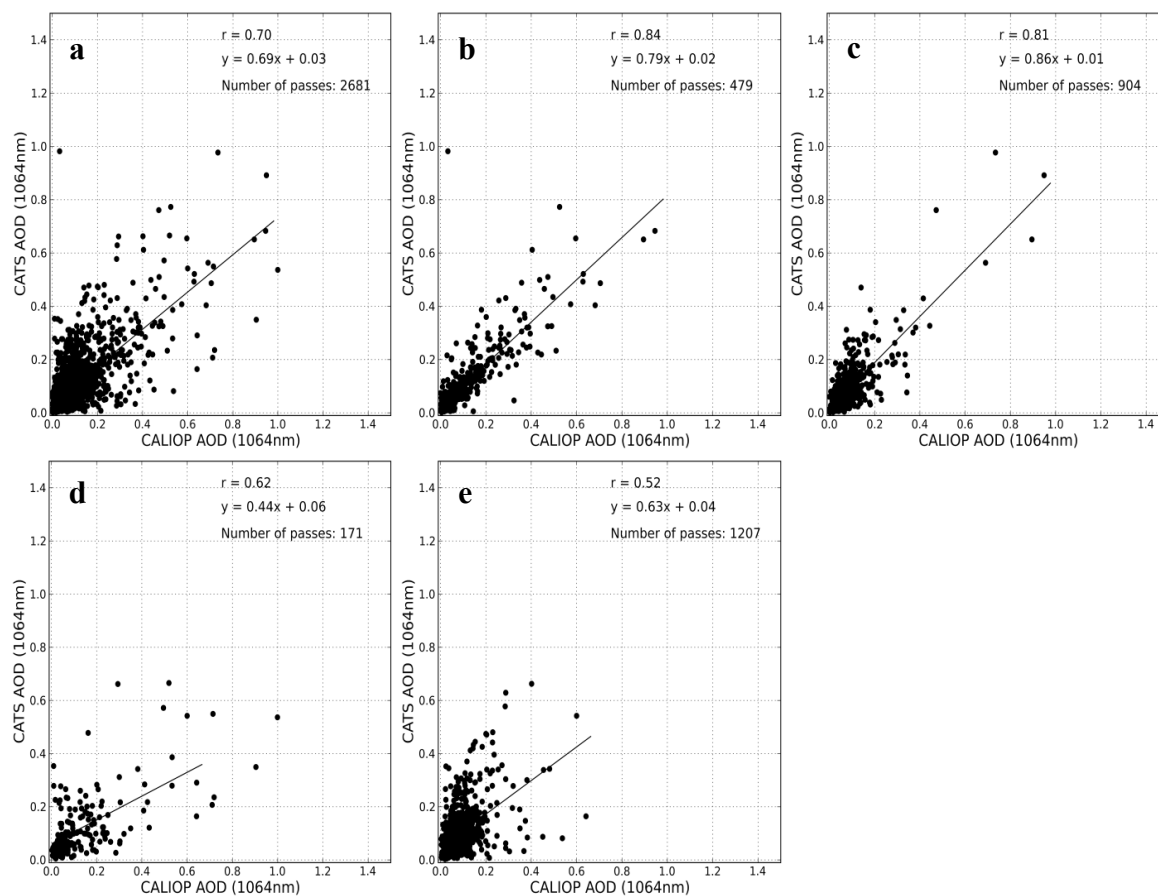


Figure 3. Collocated CALIOP 1064 nm AOD vs. CATS 1064 nm AOD with CATS QA applied for a) both day and night, b) nighttime over-land, c) nighttime over-water, d) daytime over-land, e) daytime over-water.



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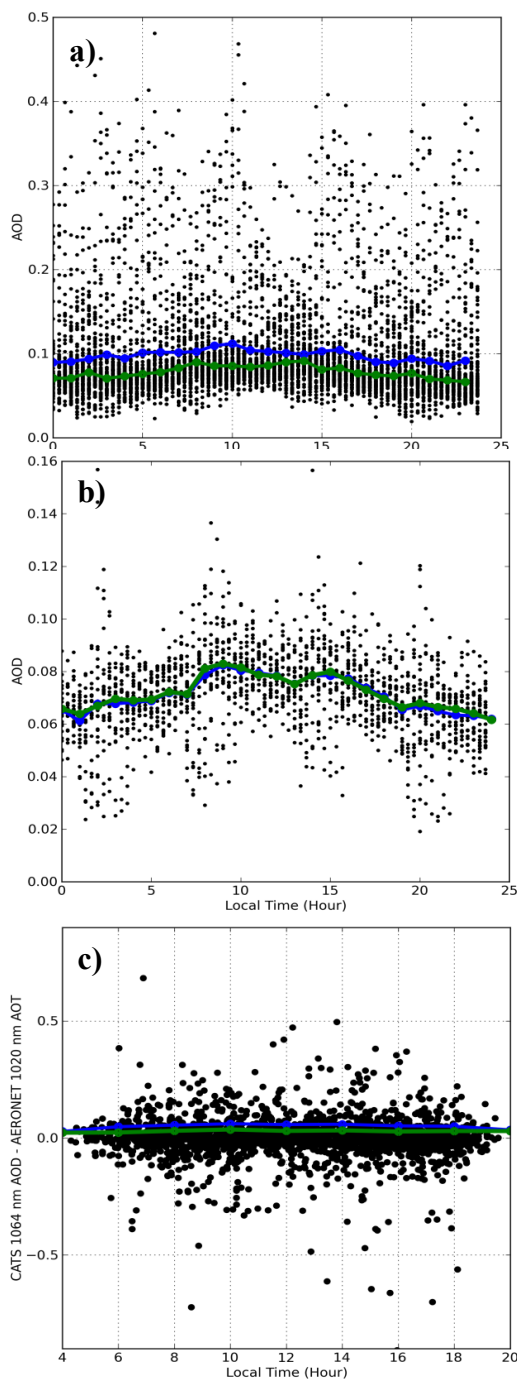


Figure 4: CATS 1064 nm AOD a) as a function of local time for the globe, and b) as a function of local time for areas south of -25 degrees. The difference between CATS 1064 nm AOD and AERONET 1020 nm AOD as a function of local time is shown in c). The mean is represented by the blue line, while the median is the green line.

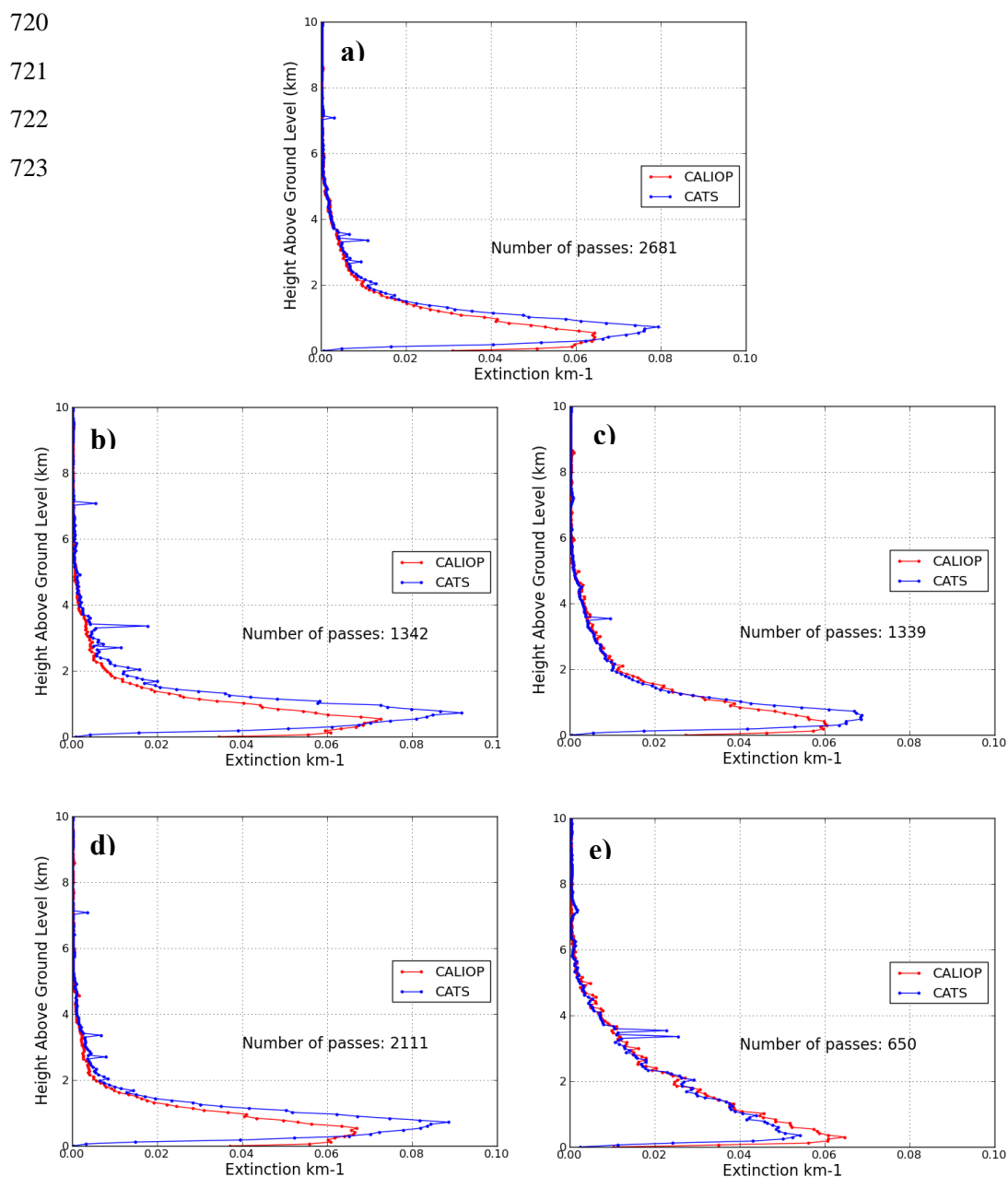


Figure 5. CATS and CALIOP vertical profiles of 1064 nm extinction for a) all profiles, b) daytime only, c) nighttime only, d) over-water, and e) over land.

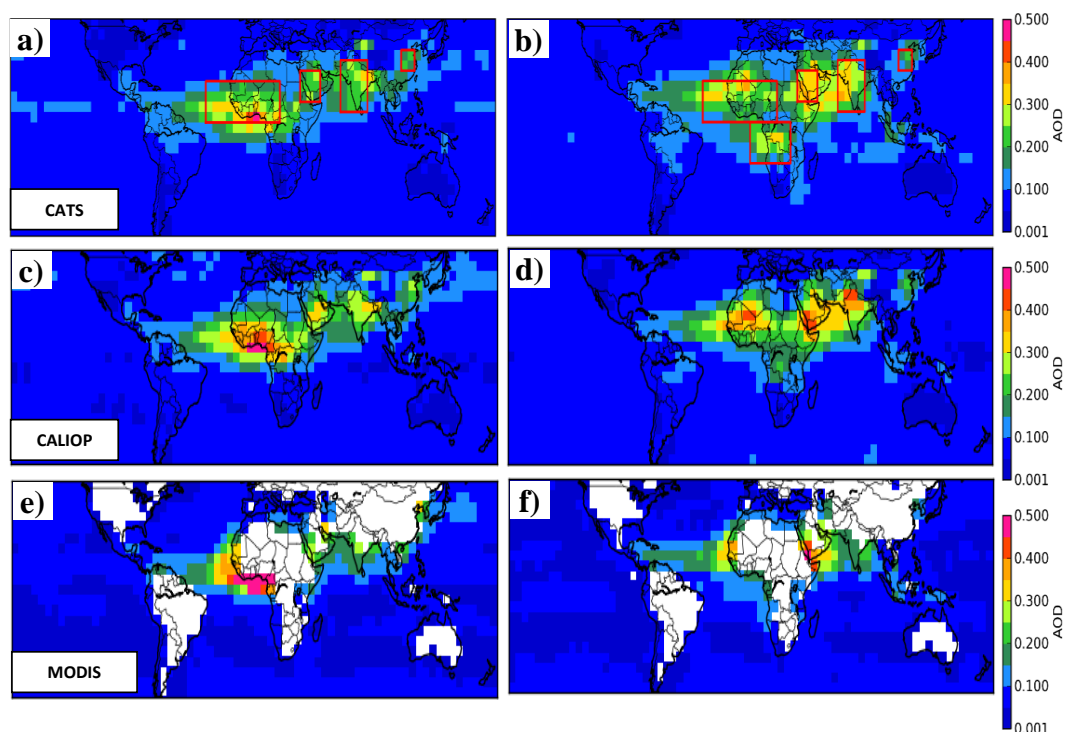


Figure 6. Mean AOD (1064 nm) by season for a) DJFMAM CATS, b) JJASON CATS, c) DJFMAM CALIOP, d) JJASON CALIOP, e) DJFMAM MODIS Aqua, and f) JJASON MODIS Aqua. Red boxes indicate locations of regional vertical distributions in Figures 12 and 13.

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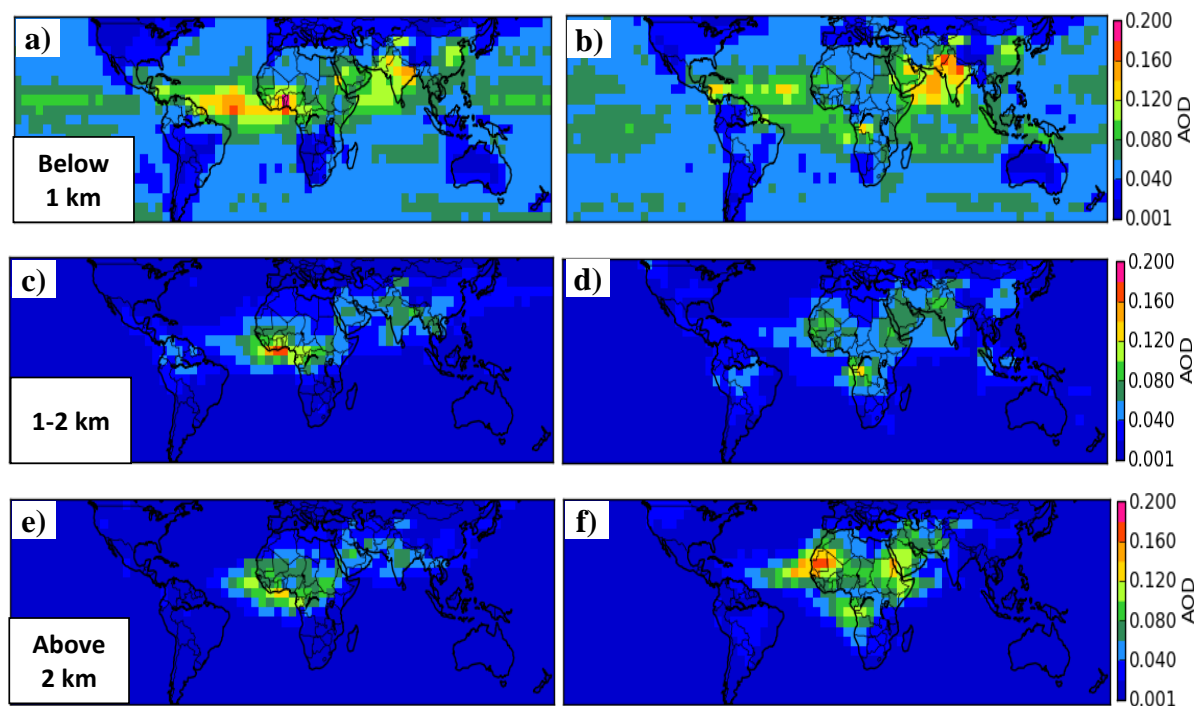


Figure 7: Mean CATS AOD (1064 nm) by season for a) DJFMAM below 1km AGL, b) JJASON below 1 km AGL, c) DJFMAM 1-2 km AGL, d) JJASON 1-2 km AGL, e) DJFMAM above 2 km AGL, and f) JJASON above 2 km AGL.



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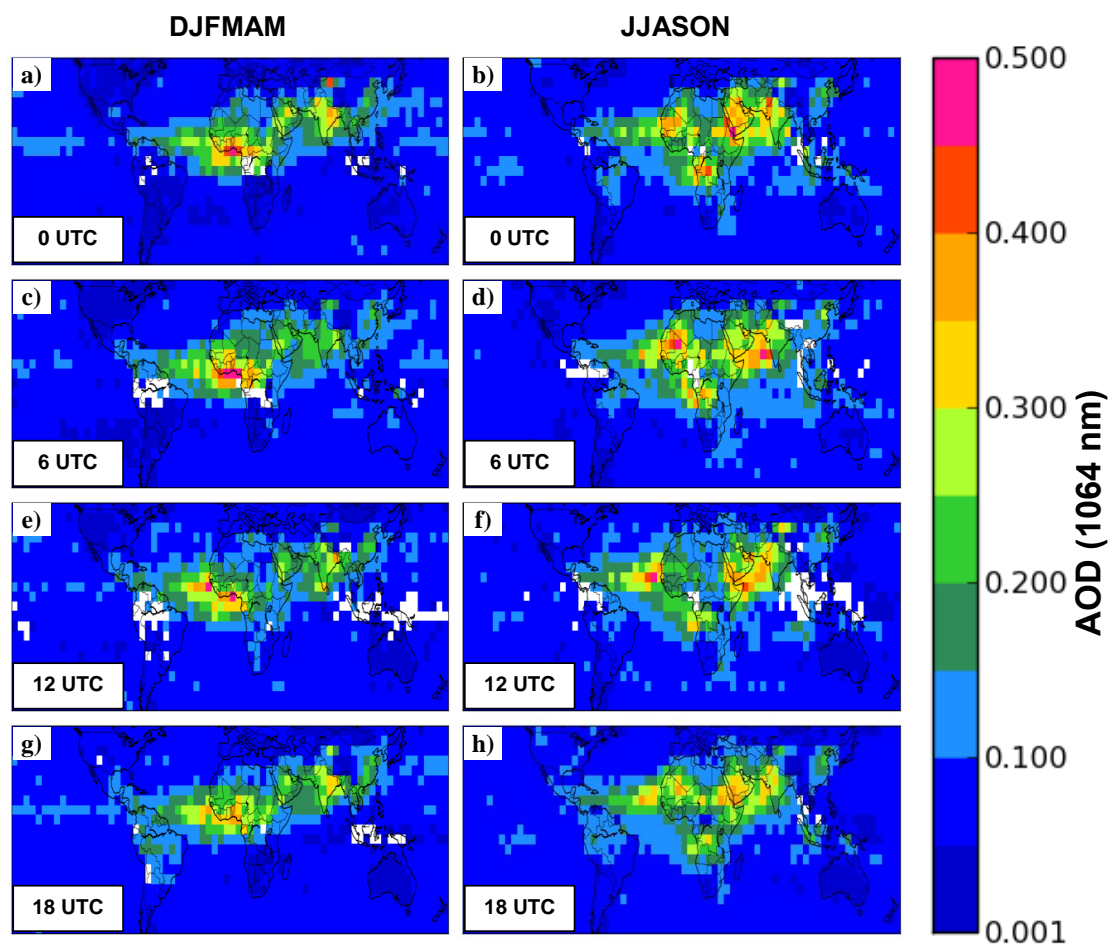


Figure 8. Seasonal Mean AOD (1064 nm) binned by every 6-hours for a) DJFMAM 0 UTC, b) JJASON 0 UTC, c) DJFMAM 6 UTC, d) JJASON 6 UTC, e) DJFMAM 12 UTC, f) JJASON 12 UTC, g) DJFMAM 18 UTC, and h) JJASON 18 UTC.



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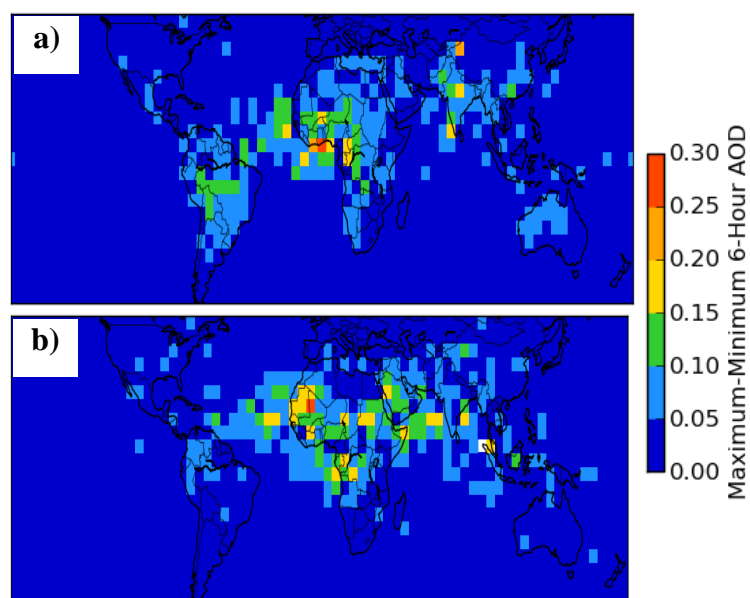


Figure 9. Maximum minus minimum mean seasonal AOD (1064 nm) for a) DJFMAM, and b) JJASON.



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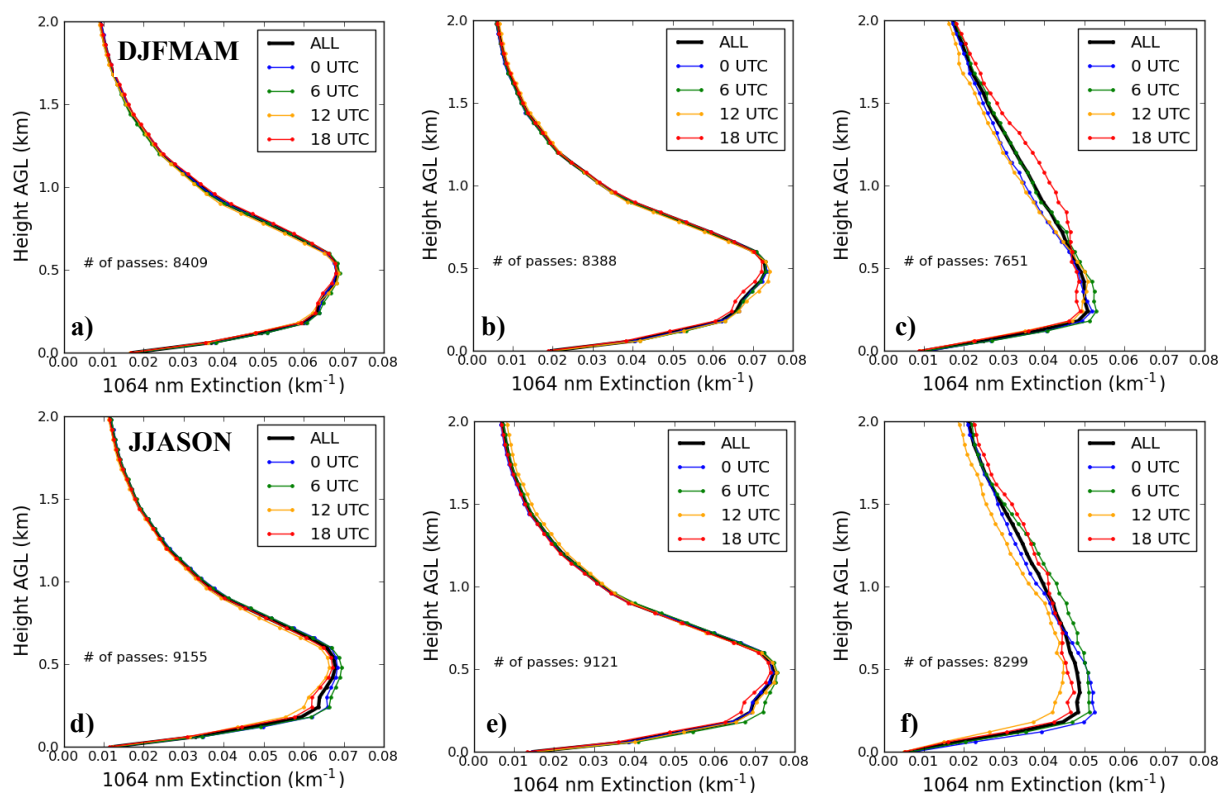


Figure 10. Global mean 6-hourly vertical profiles of CATS 1064 nm extinction for a) DJFMAM all profiles, b) DJFMAM water profiles, c) DJFMAM not-water profiles, d) JJASON all profiles, e) JJASON water profiles, f) JJASON not-water profiles.

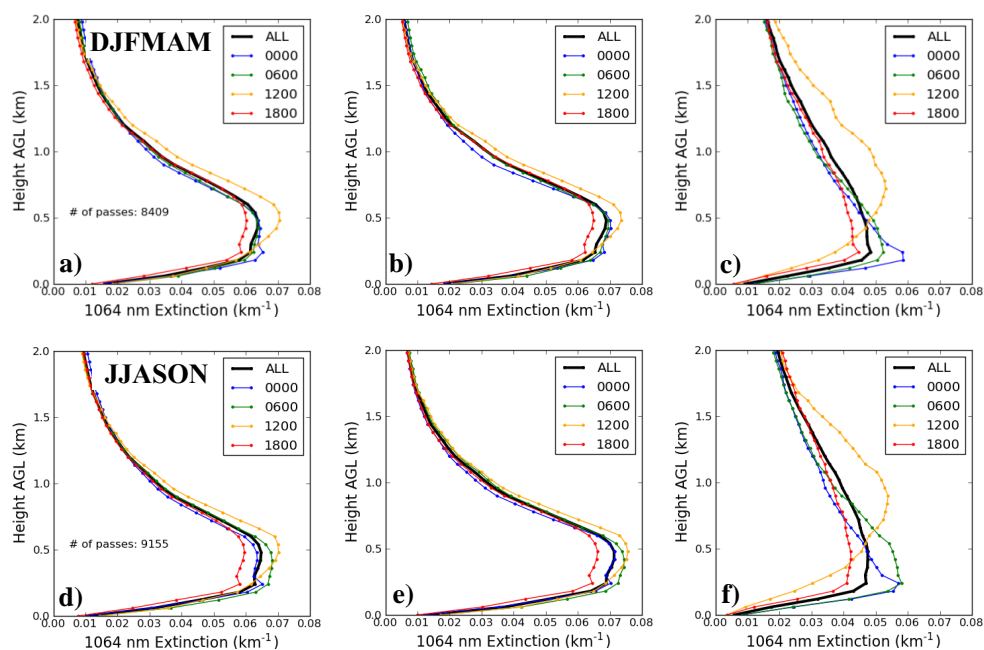


Figure 11. Global mean 6-hourly local time (0:00 am, 6:00 am, 12:00 pm and 6:00 pm) vertical profiles of CATS 1064 nm extinction for a) DJFMAM all profiles, b) DJFMAM water profiles, c) DJFMAM not-water profiles, d) JJASON all profiles, e) JJASON water profiles, f) JJASON not-water profiles.

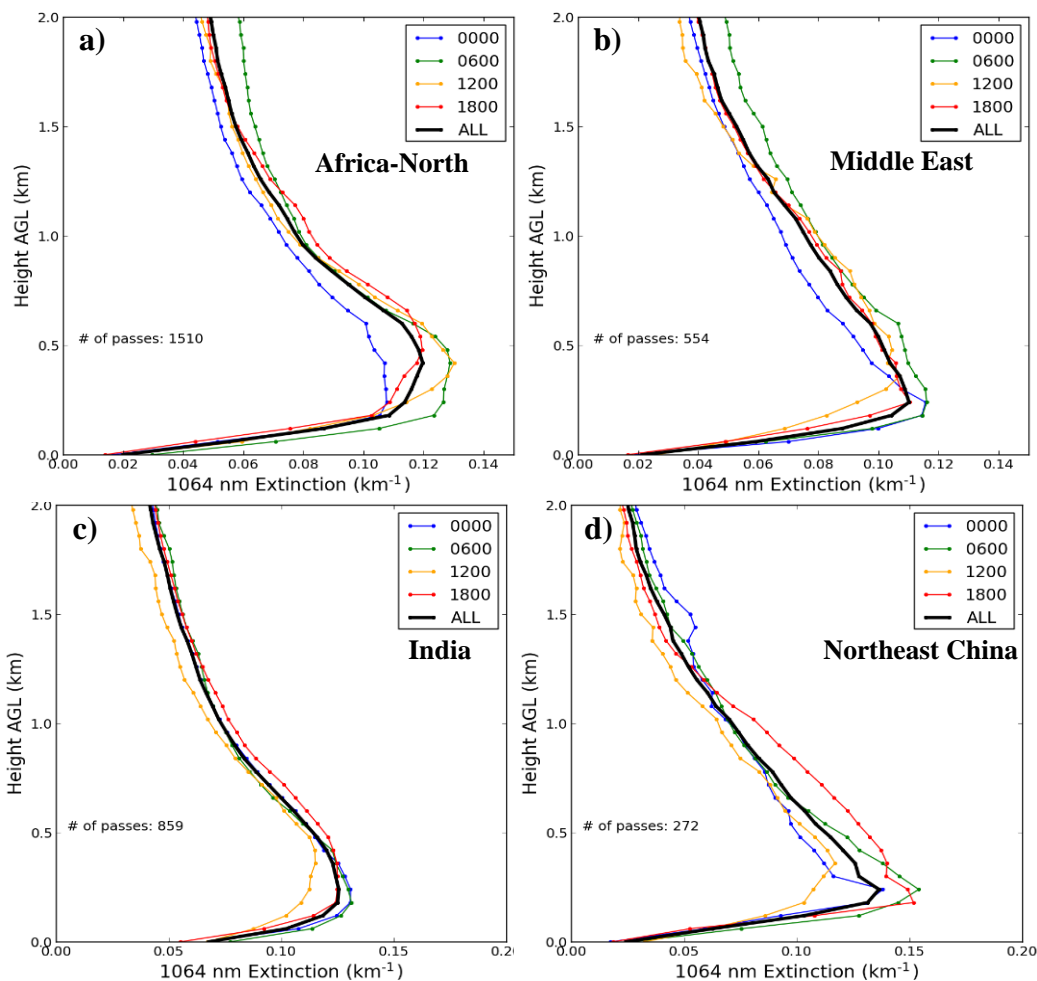


Figure 12. DJFMAM 6-hourly average (local time; 0:00 am, 6:00 am, 12:00 pm and 6:00 pm) vertical profiles of CATS 1064 nm for locations shown in Figure 6a; a) Africa-north, b) Middle East, c) India, and d) Northeast China.

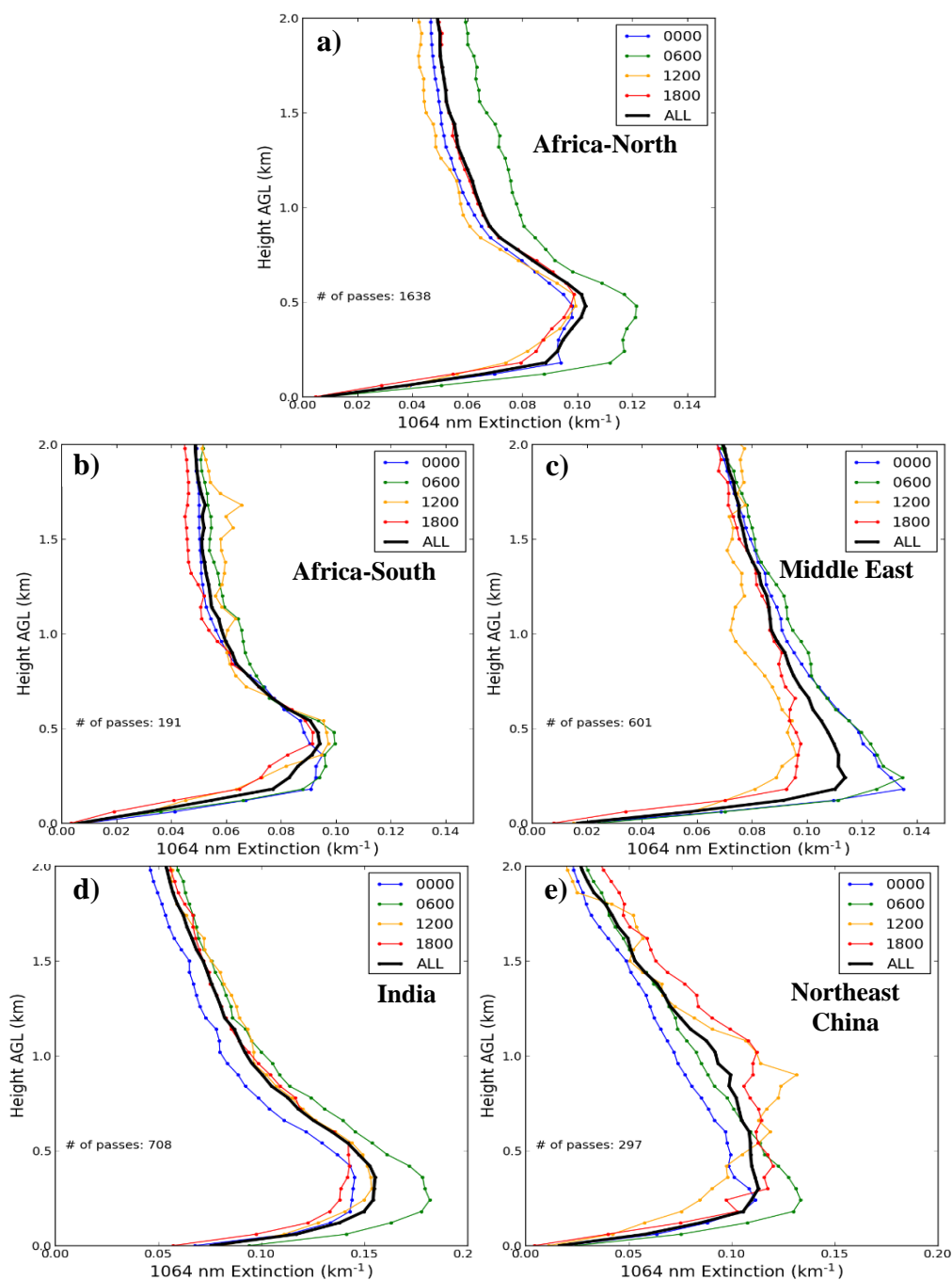


Figure 13. JJASON 6-hourly average (local time; 0:00 am, 6:00 am, 12:00 pm and 6:00 pm) vertical profiles of CATS 1064 nm for locations shown in Figure 6b; a) Africa-north, b) Africa-south, c) Middle East, d) India, and e) Northeast China.