We thank the reviewer for his/her constructive comments. Response to the Specific comments.

# About the reply to my Commnet1>

**Comment 1:** Thank you for the detailed explanation on the investigation of the spatial representativeness of the measuring stations. I think this is a reasonable approach. The only thing I wanted to ask in the previous comment was what the value on the right axis of Figure R1 in the previous response document represents. According to the response document this time, it was added by just some mistake, and I think that it was only necessary to look at the value of the left axis. If this recognition is correct, there is no further request for correction from me.

**Reply:** Thanks for your insightful comments. In the previous response documents, the right axis was added by a mistake. We are very sorry for this mistake.

# <About the other replies>

**Comment 2:** Fig.5: Please indicate whether the upper and lower figures are for summer or winter. I think it is also necessary to define the periods for spring, summer, autumn, and winter somewhere in the text.

**Reply:** We agree. In the revised manuscript, we pointed out that the upper and lower panels represent values in summer and winter, respectively. We also defined the periods for spring (March-April-May), summer (June-July-August), autumn (September-October-November), and winter (December-January-February) in the text.

**Comment 3:** Equation (A2) of Appendix A: I think "yi" should be "yij". **Reply:** We correct it.

**Comment 4:** Fig.3 and Fig.4: I think it is necessary to explain which models the gray line is pointing to.

Reply: We add it in the revised manuscript.

We thank the editor for your constructive comments. Response to the Specific comments.

**Comment 1:** This time the manuscript is much improved according to the previous comments. The reviewer #1 now provides smaller number of minor comments, which should be addressed. Despite previous request to recheck English, the Co-Editor still finds numerous errors and points of revions as follows. They are not all. Thorough checking is mandate, before recommended for publication.

**Reply:** Thanks a lot. We rechecked the English language in the revised manuscript.

(pages and lines are for the cleaned version)

**Comment 1:** 1. Abstractm line 1. Rephrase "long-term" as one-year of study is short. For example, Spatio-temporal variations of ozone (O3) and nitrogen oxides (NOx) concentrations from ...

**Reply:** We totally agree and revise it in the new manuscript.

**Comment 2:** Page 2, Line 8. Compared to MICs-Asia II with a study period of four months, the evaluation WITH observations was extended... **Reply:** We correct it.

**Comment 3:** Page 3, line 26. Page 4, line 13. extra space? **Reply:** We correct it.

**Comment 4:** Page 4, line 6. The study shows **Reply:** We correct it.

**Comment 5:** Page 4, lines 10-11. Incomplete sentence (rephrase "have rendering explaining..") **Reply:** We correct it.

**Comment 6:** Page 4, line 17. have evolved into **Reply:** We correct it.

**Comment 7:** Page 5, line 10. Use lower case for O3 **Reply:** We correct it.

**Comment 8:** Page 5, line 18. remove extra space and insert space before 2013 **Reply:** We correct it.

**Comment 9:** Page 6, line 4. Remove period character after hPa **Reply:** We correct it.

**Comment 10:** Page 6, Line 8. remove "are" **Reply:** We correct it.

**Comment 11:** Page 7, line 4. 2018 not 2017? Check all references. **Reply:** We correct it.

Comment 12:: Page 7, lines 24-26 and Page 8, lines 12-14. Similar sentences. Remove one of them Reply: We correct it.

**Comment 13:** Page 8, line 29, oxides **Reply:** We correct it.

**Comment 14:** Page 10, line 19. Insert space after O3 **Reply:** We correct it.

**Comment 15:** Page 10, line 19. Check font of "Most of the" **Reply:** We correct it.

**Comment 16:** Page 11, lines 6, 17, and 18. Model never measures. Better rephrased to "difficulty in estimating", "in simulating summer patterns", and "simulating winter patterns"

**Reply:** We correct it.

**Comment 17:** Page 12, Line 5. all models showed **Reply:** We correct it.

**Comment 18:** Page 12, Line 7. High correlations were found **Reply:** We correct it.

**Comment 19:** Page 12, Line 15, Add space after O3. **Reply:** We correct it.

**Comment 20:** Page 12, Line 17. Remove extra space after overestimation **Reply:** We correct it.

**Comment 21:** Page 12, Line 24, Remove "and" before "likely" **Reply:** We correct it.

**Comment 22:** Page 13, Line 5 and Table 4. Please double check consistency in the ranges between text and Table. Line 6. For NMBs for EA1, high values from M8-10 are likely omitted. Line 7. Correlation coefficients. Not 0.54 but 0.43, if including M8. **Reply:** Thanks. We check the statistical numbers in Table 4. We made a mistake for

NO2 NMBs of M8-M10. They should be 14.32%, 32.30% and 10.61%, we omit the character"%". In the revised manuscript, we correct these numbers.

Line 7, we correct the numbers of NMBs in EA3 in the text. "In EA3, correlation coefficients ranged from 0.43-0.72. NMBs and RMSEs except for those of M8 ranged from -0.23-0.46 and 0.90-1.79 ppbv, respectively"

**Comment 23:** Page 13, Line 9. Incomplete sentence. **Reply:** We correct it.

**Comment 24:** Page 13, Line 11. Not 0.5 but 0.43 when including M2. **Reply:** We correct it.

**Comment 25:** Page 13, Line 12. Which model has -0.42? **Reply:** We correct it. "In EA3, correlation coefficients ranged from 0.43-0.72. NMBs and RMSEs except for those of M8 ranged from -0.23-0.46 and 0.91-1.79 ppbv, respectively"

**Comment 26:** Page 13, Line 14. Change font of "the" **Reply:** We correct it.

**Comment 27:** Page 14, Line 25. The high values **Reply:** We correct it.

**Comment 28:** Page 15, line 9. Insert space after "lowest levels of" **Reply:** We correct it.

**Comment 29:** Page 16, Line 3. Lowercase for O3 **Reply:** We correct it.

**Comment 30:** Page 16, lines 10-11. Different sites for II and III for same labels? (e.g., site 4 is Sado-seki in a viewgraph and is Oki for another panel?) **Reply:** We correct it." site 4: Oki, site 5: Hedo and site 6: Banryu"

**Comment 31:** Page 16, Line 12. remove "been" **Reply:** We correct it.

**Comment 32:** Page 16, Line 16, enhanced concentrations simulated for **Reply:** We correct it.

**Comment 33:** Page 17, line 1. replace "be" with "were" **Reply:** We correct it.

**Comment 34:** Page 17, Line 2 and 5. dry deposition (not depositions) **Reply:** We correct it. **Comment 35:** Page 19, Line 15. Change ~ to – **Reply:** We correct it.

**Comment 36:** Page 19, line 27. remove extra space before Guo **Reply:** We correct it

**Comment 37:** Page 20, Line 5, add space before 2016 **Reply:** We correct it.

**Comment 38:** Page 20, line 6. Remove "and" before "likely" **Reply:** We correct it.

**Comment 39:** Page 20, line 1 and 11. Add space before 1996 and before 100-200 m **Reply:** We correct it.

**Comment 40:** Page 20, line 15. were not was **Reply:** We correct it.

**Comment 41:** Page 20 Line 19. East Asia **Reply:** We correct it.

**Comment 42:** Page 28, Remove ACPD for Li, M. et al. **Reply:** We correct it.

# Model evaluation and inter-comparison of surface-level ozone and relevant species in East Asia in the context of MICS-Asia phase III Part I: overview

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<sup>18</sup>Beijing Normal University, Beijing, 100875, China

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Abstract: Long-termSpatio-temporal variations of ozone (O<sub>3</sub>) and nitrogen oxide (NO<sub>x</sub>) from fourteen state-of-the-art chemical transport models (CTMs) are intercompared and evaluated with O<sub>3</sub> observations in East Asia, within the framework of the Model Inter-Comparison Study for Asia phase III (MICS-ASIA III), designed to evaluate the capabilities and uncertainties of current CTMs

- 5 simulations for Asia and to provide multi-model estimates of pollutant distributions. These models were run by fourteen independent groups working in China, Japan, South Korea, the United States and other countries/regions. Compared to MICS-Asia II, the evaluation of with observations was extended to one-full year across China and the western Pacific Rim from four months. In general, model performance levels for O<sub>3</sub> varied widely by region and season. Most models captured key patterns of
- 10 monthly and diurnal variation of surface O<sub>3</sub> and its precursors in North China Plain and western Pacific Rim but failed to do so for the Pearl River Delta. A significant overestimation of surface O<sub>3</sub> was evident from May-September/October and from January-May over the North China Plain, western Pacific Rim and Pearl River Delta. Comparisons drawn from observations show <u>that the</u> considerable diversity in photochemical production partly contributed to this overestimation and to high levels of
- 15 intermodel variability in O<sub>3</sub> for North China. In terms of O<sub>3</sub> soundings, the ensemble average of the models reproduced the vertical structure for the western Pacific, but overestimated O<sub>3</sub> levels to below 800 hPa in the summer. In the industrialized Pearl River Delta, the ensemble average presented an overestimation for the lower troposphere and an underestimation for the middle troposphere. The ensemble average of 13 models for O<sub>3</sub> did not always exhibit superior performance compared to certain
- 20 individual models in contrast to its superior value for Europe. This finding suggests that the spread of ensemble-model values does not represent all uncertainties of O<sub>3</sub> or that most MICS-Asia III models miss key processes. This study improves the performance of modeling O<sub>3</sub> in March at Japanese sites than the previous phase of MICS-Asia (MICS-Asia II). However, it overpredicts surface O<sub>3</sub> concentrations for western Japan in July, which was not found by MICS-Asia II. Major challenges still
- remain in regard to identifying the sources of bias in surface O<sub>3</sub> over East Asia in CTMs.

#### 1. Introduction:

Tropospheric ozone (O<sub>3</sub>) is a significant secondary air pollutant produced through thousands of photochemical reactions and that is detrimental to human health, ecosystems, and climate change as a strong oxidant (WHO, 2005; The Royal Society, 2008). With rapid industrialization and urbanization in

- 5 the last two decades,  $O_3$  concentration is rising at a higher rate in East Asia than in other regions, and on 30% of days in megacities (e.g. Beijing, Shanghai Guangzhou in China) values exceed air quality standard of World Health Organization (100  $\mu$ g/m<sup>3</sup>) for 8-hour average surface  $O_3$  concentration (Wang et al., 2017). High  $O_3$  concentrations have received more attention from the public and from policymakers in East Asia. The Ministry of Environment Japan has imposed stringent measures to reduce
- traffic emissions since the 1990s, and non-methane volatile organic compounds (NMVOCs) and NO<sub>x</sub> mixing ratios have decreased by 40-50% and 51-54%, respectively (Akimoto et al., 2015). In 2012, China released a new ambient air quality standard under which a limit on the 8-hour O<sub>3</sub> maximum was set for the first time. However, these measures do not prevent the persistent increase of the ground-level O<sub>3</sub> in East Asia. The average mixing ratio of O<sub>3</sub> has increased 20-30% in Japan over the last 20 years (Akimoto et al., 2015). In Chinese megacities, 8-hr O<sub>3</sub> concentrations have increased 10-30%

since 2013 (Wang et al., 2017).

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The main method used for the detailed evaluation of effects of air quality policies at the scale of East Asia is that of numerical air quality modeling. Several global and regional scale CTMs (e.g. GEOS-Chem, CHASER, CMAQ, CAMx, WRF-Chem and NAQPMS) have been developed over the past few decades and have been widely used to simulate the O<sub>3</sub> formation process and to evaluate strategies for its control (Streets et al., 20087; Li et al., 2007; 2008; Yamaji et al., 2006; – Zhang et al., 2008; Liu et al., 2010; Wang et al., 2013; He et al., 2017; Nagashima et al., 20170). Such simulations have identified the key precursors of O<sub>3</sub> formation in East Asia (Zhang et al., 2008; Liu et al., 2010; Tang et al., 20140; He et al., 2017), have assessed the contributions of international and regional

25 transport (Streets et al., 2008; Li et al., 2008), and have predicted O<sub>3</sub> mixing ratios under different future emission scenarios (Wang et al., 2013). However, discrepancies remain between models and observations<sub>a5</sub> indicating that model simulations of O<sub>3</sub> in East Asia still need to be improved (Han et al., 2008). Modeling uncertainties related to emissions, chemistry, wet and dry deposition, and transport

can hardly be addressed using a single model. Model inter-comparison has thus been recognized as an effective way to address problems and has been successfully applied in Europe and North America in phase 2 of the Air Quality Model Evaluation International Initiative (AQME II; Rao et al., 2011). Limited model inter-comparison related to air quality in East Asia has been conducted. Phases I and II

- 5 of the Model Inter-Comparison Study for Asia (MICS-Asia) were initiated in 1998 and 2003, to explore the potential sources of model uncertainties regarding sulfur, O<sub>3</sub>, nitrogen compounds and aerosols (Carmichael et al., 2002, 2008). They study shows that the predicted temporal variations found of for surface O<sub>3</sub> in eight regional CTMs generally tend to be lower than those observed in 2001 with poor correlations in the western Pacific in March and December (Han et al., 2008). Model performance
- 10 levels for O<sub>3</sub> were found to vary greatly in southern China. Inconsistencies inf horizontal grids, emissions and meteorological inputs used among models increased have rendering rendered explaining intermodel variability using in MICS-Asia II results more difficult. More importantly, model evaluations for industrialized China have not been conducted due to a lack of observations, which has been detrimental to efforts made to improve O<sub>3</sub> model performance levels .levels.
- 15 Recently, regional CTMs have been greatly improved by coupling more mechanisms (e.g. heterogeneous chemistry and on-line calculation of photolysis rates) and accurate chemical reaction rates. For example, gas-phase chemistry mechanisms of Models 3-Community Multiscale Air Quality (CMAQ) have been <u>developed-evolved</u> into CBM05 and SAPRC07 from CB04 and SAPRC99. It is critical to evaluate the updated models' abilities to simulate current air quality levels over East Asia. In
- 20 2010, MICS-Asia was expanded to Phase III wherein, 13 regional CTMs and 1 global CTM were run over one-full year by 14 independent groups from East Asia and North America, using a common reference model input data-set (namely, the emission inventory, meteorological fields and horizontal grids). In addition to observations made in Japan by the Acid Deposition Monitoring Network in East Asia (EANET) that were used for MICS-Asia II, new observational data from China were made
- 25 available for MICS-Asia III and were obtained from the Chinese Ecosystem Research Network (CERN) and the Pearl River Delta Regional Air Quality Monitoring Network (PRD RAQMN). An intercomparison of CTMs in China, Japan and the western Pacific for one full year had never before

been performed, creating a broader database to use for comparisons. The completeness of MICS-Asia III is therefore unique.

In this paper, we mainly evaluate the capacities of participating models in MICS-Asia III to simulate concentrations of O<sub>3</sub> and its related species within the MICS-Asia III framework. The

5 follwingfollowing questions are addressed: (1) How well do various air quality models perform in simulating O<sub>3</sub> levels in East Asia? (2) How consistent or discrepant are the models? (3) How do mutimodel ensembles improve O<sub>3</sub> simulation accuracy? This paper is expected to provide valuable insights into the capacities and limitations of CTMs when applied to East Asia.

# 2. Models and data

# 10 2.1 Experimental set up

In this study, all participating models were run for the year 2010 and provide<u>d</u> gridded monthly mean diurnal <u>mixing ratios of</u>  $O_3$  and its precursors <u>mixing ratios</u> in the lowest model layer. For  $O_3$ , monthly three-dimensional data were also submitted.

# 2.2 Participating models and input data

intercomparison projects (e.g., AQME II).

Table 1 summarizes the specifications of participating CTMs. These models include two versions of CMAQ (v4.7.1 and 5.0.2; Byun and Schere, 2006), the Weather Research and Forecasting model coupled with Chemistry (WRF-Chem; http://www.acd.ucar.edu/wrf-chem), the Nested Air Quality Prediction Modeling System (NAQPMS; Li et al., 2007), the Japan Meteorological Agency (JMA)'s non-hydrostatic meteorology-chemistry model (NHM-Chem; Kajino et al., 2012), the NASA-Unified
Weather Research and Forecasting (NU-WRF; Tao et al., 2013) and GEOS-Chem (http://acmg.seas.harvard.edu/geos/). They have been documented in the scientific literature and have been widely applied in modeling studies of East Asia. Table 1 does not list model names to maintain

each model's anonymity. Similar behavior was observed from MICS-Asia II and other model

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MICS-Asia III participants were provided with a reference meteorological field for the year 2010, generated with the Weather Research and Forecasting Model (WRF) version 3.4.1 model. The domain

of meteorological fields is shown in Fig. 1. WRF v3.4.1 are-is\_driven by the final analyses dataset (ds083.2) from the National Centers for Environmental Prediction (NCEP), with  $1^{\circ} \times 1^{\circ}$  resolution and a temporal resolution of 6 h. A four-dimensional data assimilation nudging toward the NCEP dataset was performed to increase the accuracy of the WRF. The horizontal model domain of 182 ×172 grids

- on a Lambert conformal map projection with 45-km horizontal resolution, is shown in Fig. 1.
   Vertically, the WRF grid structure consists of 40 layers from the surface to the top of model (10 hPa-).
   Standard meteorological fields were applied by the majority of groups. Several other models were employed to perform simulations using their own meteorological models (e.g., RAMS-CMAQ and GEOS-Chem). The WRF-Chem utilized the same model (WRF) as the standard meteorological
- 10 simulation but considered the feedback of pollutants to meteorological fields. Consequently, their meteorological fields are-may be slightly different from the standard. GEOS-Chem is driven by the GEOS-5 assimilated meteorological fields taken from the Goddard Earth Observing System of the NASA Global Modeling Assimilation Office. The couples of meteorological data and CTMs variedy for each group, likely resulting in a diversified set of model outputs.
- 15 MICS-Asia III provides a set of monthly anthropogenic emission inventories for the year 2010 called MIX (Li et al., 2016). MIX is a mosaic of up-to-date regional and national emission inventories that includes Regional Emission inventory in ASia (REAS) version 2.1 for the whole Asian region (Kurokawa et al., 2013), the Multi-resolution Emission Inventory for China (MEIC) developed by Tsinghua University, a high-resolution NH<sub>3</sub> emission inventory by Peking University (Huang et al.,
- 20 2012), an Indian emission inventory developed by Argonne National Laboratory (ANL-India, Lu et al., 2-011; Lu and Streets, 2012), and the official Korean emission inventory from the Clean Air Policy Support System (CAPSS; Lee et al., 2011). Biogenic emissions are taken from the Model of Emissions of Gases and Aerosols from Nature (MEGAN). Hourly biogenic emissions were obtained for the entire year of 2010 using version 2.04 (Guenther et al., 2006). Biomass burning emissions were processed by
- 25 regridding Global Fire Emissions Database version 3 (GFEDv3) (0.5 by 0.5 degree). Volcano SO<sub>2</sub> emissions were provided, with a daily temporal resolution by the Asia Center for Air Pollution Research (ACAP). The MICS-ASIA III emission group directly prepared a gridded inventory according to the configuration of each CTM. NMVOC emissions are were spectated into model-ready

inputs for three chemical mechanisms: CBMZ, CB05 and SAPRC-99. Weekly and diurnal profiles were also provided. The standard emission inventory was applied by all models. The majority of models employed official suggested vertical and time profiles of pollutants from each sector by emission group. M13 and M14 make the projections by themselves. More information can be found in

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Li et al. (2017) and Gao et al. (20187).

MICS-Asia III also provided two sets of chemical concentrations for the top and lateral boundaries of the model domain, which were derived from 3-hourly global model outputs for the year 2010. The global models were run by University of Tennessee (http://acmg.seas.harvard.edu/geos/) and Nagoya University (Sudo et al., 2002). GEOS-Chem was run with a 2.5°×2° horizontal resolution and

10 47 vertical layers by University of Tennessee, and Chemical AGCM for Study of Atmospheric Environment and Radiative Forcing (CHASER) was run with a 2.8°× 2.8° horizontal resolution with 32 vertical layers by Nagoya University. Some models applied boundary conditions depending on their own past experiences.

# 2.3 Observational data for O<sub>3</sub>

climate.

15 In this study, East Asia was divided into three subregions as shown in Fig. 1. The selection of subregions was based on emission, climate and observation data coverage. The North China Plain (EA1) and Pearl River Delta (EA2) represent highly industrialized regions of the mid-latitudes. EA1 is characterized by a temperate and tropical continental monsoon climate with marked seasonality. EA2 is located in southern China and is less affected by continental air masses. EA3 covers the northwest 20 Pacific and the Sea of Japan and represents the downwind regions of the Asian continent with a marine

Hourly O<sub>3</sub> and NO<sub>x</sub> observations for the year 2010 in East Asia were obtained from the CERN, PRD-RAQMN, and EANET. The CERN was built by the Institute of Atmospheric Physics, Chinese Academy of Sciences and includes 19 surface stations covering an area of  $500 \times 500$  km<sup>2</sup> across North

25 China Plain (EA1 sub-region; Ji et al., 2012). The stations were set up according to United States Environmental Protection Agency method designations. Half of them were remote, rural, suburban and clear urban sites. Nine sites were located within meteorological stations or on campuses of universities in urban regions, with little influence from local sources and sinks. A comparison of NO and ethene-

emission rates at these sites in 45 km and 3 km resolution emission inventories shows that observations generally represent the - 45 km averages of pollutants. The PRD RAQMN was jointly established by the governments of Guangdong Province and the Hong Kong Special Administrative Region and consists of 16 automatic air quality monitoring stations located across the EA2 subregion (Zhong et al.,

- 5 2013). Thirteen of these stations are operated by the Environmental Monitoring Centers in Guangdong Province and the other three are located in Hong Kong and are managed by the Hong Kong Environmental Pollution Department. The PRD RAQMN was designed to probe regional air quality, to assess the effectiveness of emission reduction measures and to enhance the roles of monitoring networks in characterizing regional air quality and in supporting air quality management. Thus, the
- sites are rarely influenced by local sources and sinks. The EANET was launched in 1998 to address acid deposition problems in East Asia, following the model of the Cooperative Program for Monitoring and Evaluation of the Long-Range Transmission of Air Pollutants in Europe. In this study, eight remote stations in the northwestern Pacific and Japan (EA3 subregion) were selected to evaluate model performance levels for the downwind regions of the Asian continent (Ban et al., 2016). More
- 15 information on the EANET can be found at http://www.eanet.asia/. Note that only stations with at least 75% data validity were chosen. Table S1 in the supplementary section provides detailed site description. Our comparisons of NO<sub>x</sub> and VOCs emission rates conducted on grids for these stations at 45 km and 3 km resolution emission inventories suggest that our selected stations have rarely received local emissions.
- 20 O<sub>3</sub> wasere measured by Thermo Scientific 49i with UV photometric technology in CERN network and by Thermo Scientific 49C in PRD-RAQMN and EANET network. NO<sub>x</sub> was measured by Thermo Scientific 42C NO-NO<sub>2</sub>-NO<sub>x</sub> Analyzer with chemiluminescence technology at 40 sites in all three networks (CERN, PRD-RAQMN and EANET). NO<sub>x</sub> measurements exhibited sometimes biases (especially for stations located far from sources) when using molybdenum converter devices since all nitrogen oxiydes were measured. This bias was found to be dependent on chemical conditions. A onemonth continuous measurement collected in August with a chemiluminescence analyzer and Aerodyne
  - from the chemiluminescence analyzer is minor when NO2 concentrations exceed 10-15 ppbv, ranging

Cavity Attenuated Phase Shift Spectroscopy (CAPS) from an urban site in Beijing shows that this bias

from 10% to 30% under low NO<sub>2</sub> conditions (<10 ppbv) (Ge et al., 2013). Measurements collected from a rural site in South Korea reveal a similar pattern across all seasons (Jung et al., 2017). These comparisons suggest that observations made using molybdenum converters may overestimate NO<sub>2</sub> by 10-20% for EA1 and EA2 and 30% for EA3, introducing uncertainties into the NO<sub>2</sub> model evaluation in this study.

#### 3. Model validation and general statistics

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#### 3.1 Annual concentrations of surface O<sub>3</sub>, nitric oxide (NO) and nitrogen dioxide (NO<sub>2</sub>)

Fig. 2 provides a concise comparison of model performances for annual O<sub>3</sub>, NO and NO<sub>2</sub> for three sub regions of East Asia. A box-and-whisker representation was used to show the frequency

- distribution of monthly concentrations measured from stations in each subregion. The O<sub>3</sub> normalized mean bias (NMB) and root mean square error (RMSE) of the ensemble mean were found to be significantly less than the ensemble median in most cases (Table 2). Therefore, we only present multi-model mean ensemble results (Ense). In general, the majority of the models significantly overestimated annual surface O<sub>3</sub> relative to observations oinf EA1, EA2 and EA3. Ense overestimated surface O<sub>3</sub> by
   10-30 parts per billion by volume (ppbv) for these subregions. Ense NO<sub>2</sub> levels closely reflect
- observations to within  $\pm 20\%$  across all subregions. In EA1 and EA2, Ense NO levels were found to be

5-10 ppbv lower than those observed observations while exhibiting reasonable levels for EA3.

Of the models, M11 for subregions EA1 and EA2 and M7 in for EA2 and EA3 more closely reflect O<sub>3</sub> observations. M11 simulated O<sub>3</sub> with RMSEs of 9.5 ppbv and 13.3 ppbv for EA1 and EA2,

- 20 respectively (Table 2). The models' performance in simulating O<sub>3</sub> was found to be closely related to their performance for NO<sub>2</sub> and NO. In highly polluted regions (EA1 and EA2), a persistent underestimation of NO was evident across most models. As an interesting phenomenon, we found the models' performance regarding O<sub>3</sub> measurements to vary greatly for EA3, though M8 exhibited a consistent performance with respect to NO and NO<sub>2</sub>. This finding suggests that O<sub>3</sub> was significantly
- 25 affected by other factors in addition to local chemistry in EA3. M8 underestimated O<sub>3</sub> and overestimated NO in all subregions by 40-50%. The highest O<sub>3</sub> titration level observed in M8 may have generate lower O<sub>3</sub> levels than those indicated by other models and observations.

#### 3.2 Monthly variations of surface O<sub>3</sub>, NO and NO<sub>2</sub>

Fig. 3 presents monthly mean concentrations of O<sub>3</sub>, NO and NO<sub>2</sub> for the three subregions across East Asia. When two or more observation sites are located in the same model grid, their mean values are used to evaluate model performance. All models captured the observed seasonal cycles of O<sub>3</sub>, NO

- and NO<sub>2</sub> for EA1. From May-September, Ense O<sub>3</sub> was 10-30 ppbv higher than observed values ( 30-70% of observed values) while Ense NO and NO<sub>2</sub> levels appeared to be consistent with observations with mean biases of  $\leq 4$  3 ppbv. This finding suggests that an intercomparison of O<sub>3</sub> production efficiency levels per NO<sub>x</sub> with observations is needed. For EA2, Ense O<sub>3</sub> values agree well with observed high autumn O<sub>3</sub> levels but are overestimated from January to September by 5-15 ppbv (15-
- 10 60% of observations). This overestimation reached the highest point from in March-April (15\_ppbv) and led to a spring peak in simulated O<sub>3</sub> values not found in the observations. This overestimation is partly related to the underestimation of NO in the same months, which decreased the titration effect. For NO<sub>2</sub>, Ense value agree<u>d</u> well with observed values for June-December, and slightly underestimated observations for January-May. For EA3, the ensemble NO<sub>2</sub> was generally close to observed values
- within ±0.5 ppbv. Significant overestimations of O<sub>3</sub> and underestimations of NO were observed from June-October. Similar results have been found from MICS-Asia II and through another model inter-comparison project of the Task Force on Hemispheric Transport of Air Pollution (TF HTAP), suggesting that such results may stem from differences in representations of southwesterly clean marine air masses dispersion observed across different metrological fields used in CTMs (Han et al.,\_2009).

For individual models, M11 achieved the highest degree of model reproductivity for monthly mean O<sub>3</sub> levels in EA1-of the examined models. Most of the other models overestimated O<sub>3</sub> by 100-200% for May-October. The largest levels of model bias and intermodel variability for NO and NO<sub>2</sub> appeared in the winter, and likely came from the NO<sub>x</sub> vertical diffusion and heterogeneous chemistry

25 (Akimoto et al., 2019). In EA2, M7 seems to have achieved the highest levels of O<sub>3</sub> reproducibility. Most of the models (except for M11 and M12) show high O<sub>3</sub> concentrations for March-May and September-November. Observed O<sub>3</sub> values show that the highest concentrations appeared from October-November. M11 captured the observed January-May O<sub>3</sub> value due to relatively high NO

concentrations. However, NO was overestimated by M11 from May-September, leading to an underestimation of O<sub>3</sub> levels. In EA3, spatially averaged O<sub>3</sub> concentrations often differ by more than 20 ppbv in individual models. The highest levels of intermodel variability in O<sub>3</sub> values appeared from May-October, overestimating O<sub>3</sub> levels relative to observations by 10-40 ppbv. Interestingly, although

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M8, M9 and M14 exhibited similar magnitudes with observations for June-September, they significantly underestimated observations in other months by 200-300%. A detailed investigation is required in future studies.

#### 3.3 Diurnal concentrations of surface O<sub>3</sub>

Subregional O<sub>3</sub> diurnal variations are shown in Fig. 4. In general, model results for three
subregions exhibited a larger spread with a magnitude of 10-50 ppbv across the diurnal cycle than those observed in Europe and North America (Solazzo et al., 2012). Summer Ense O<sub>3</sub> levels exhibited a systematic pattern of overestimation (20 ppbv) throughout the diurnal cycle in EA1. This indicates that the models had difficulty <u>in measuring estimating</u> summer O<sub>3</sub> levels for the North China Plain.
Compared to summer conditions, only a slightly systematic overestimation of Ense O<sub>3</sub> levels was observed for the other seasons (3-5 ppbv). In EA2, Ense O<sub>3</sub> levels generally agree<u>d</u> with summer, autumn and winter observations. In particular, the O<sub>3</sub> maximum occurring at around noon was reasonably reproduced. Only a 3-5 ppbv overestimation was observed from\_16:00-23:00 and in early morning (6:00-10:00). In the spring, a systematic overestimation of Ense O<sub>3</sub> values was observed

across the whole diurnal cycle (5-10 ppbv). In EA3, Ense captured the minor diurnal variations in O<sub>3</sub>

20 across all four seasons, but significantly overestimated observations for the summer and autumn (5-20 ppbv). In the spring and winter, differences between Ense and observations fell within 5 ppbv.

Of all of the models, M11 exhibited the best model performance level in measuring peak daily O<sub>3</sub> concentrations of 60 ppbv from 14:00-16:00 in EA1, but it still overestimated nighttime O<sub>3</sub> levels by 10 ppbv. Compared to their performance in measuring simulating summer patterns, the models performed significantly better in measuring-simulating winter conditions due to the weak intensity of photochemical reactions except in <u>the</u> case of M2, M10 and M8. Differences between observations and most simulations for both the nighttime and daytime fell within 5 ppbv. These differences in the models' performances between the summer and winter imply that the variety of chemistry

parametrizations applied to different models partly explain the intermodel variability of simulated O<sub>3</sub> levels in EA1 (North China Plain). For EA2, the majority of models agree<u>d</u> well with diurnal variations occurring in the summer and autumn. However, most models exhibited a tendency to overestimate the O<sub>3</sub> concentrations for both the daytime and nighttime in the spring. The overestimated magnitudes

- 5 exceeded 10 ppbv and 25 ppbv (of observed values of 20-35 ppbv) for the nighttime and daytime, respectively. M11 reproduced observed O<sub>3</sub> levels for the spring but underestimated O<sub>3</sub> levels for the summer and autumn. For EA3, significant levels of intermodel variability persisted throughout the year. Amplitudes of intermodel variability except for those of M8 and M14 reached approximately 20 ppbv and 10 ppbv in the spring-summer and autumn-winter, respectively. M8 and M14 generated the lowest O<sub>3</sub> values of the models for the whole year.

# 3.4 Error statistics on surface concentrations

In this section, we present statistics on the models' performance based on monthly values. Values are calculated with equations shown in Appendix A. On a yearly basis, all models <u>showed observe</u> the highest (0.8-0.9) and lowest (0.1-0.6) correlation coefficients for O<sub>3</sub> for EA1 and EA2, respectively

(Table 2). High correlations were observed found in EA1 mainly because the summer-maximum and winter-minimum seasonal cycles are typical of polluted regions represented in all of the participating models. In general, Ense performed better than the individual models in representing NO<sub>2</sub> for East Asia, reproducing the observed seasonal cycles and magnitudes. However, Ense did not always exhibit a superior performance in measuring simulating O<sub>3</sub> levels over individual models for East Asia, which
stands in contrast to its performance for Europe (Table 2). M7 and M11 agreed well with observations for EA1 and EA2 while ENSE tendeds to overestimate O<sub>3</sub> concentrations for May-September in EA1 and for January-September in EA2. Loon et al. (2007) indicated that Ense exhibits a superior performance level only when the spread of ensemble-model values is representative of O<sub>3</sub> uncertainty. This indicates that most models do not reflect this uncertainty or missed key processes of MICS-Asia
III.

Considerable overestimation-s made by most of the models for May-September led to high NMB (0.25-1.25) and RMSE (10-33 ppbv) values for EA1. M11 generated the lowest NMB (0.09) and RMSE (9.46 ppbv) values of the examined models. For EA2, M9 and M10 generated stronger

correlations than the other models. However, their corresponding NMB and RMSE values were also the highest. These findings imply that systematic model biases are present in these two models. M7 exhibited lower NMB and RMSE values than the other models, but its correlation was measurementsmeasured as only 0.29. For EA3, correlations exhibited the largest degree of intermodel variability across all subregions, ranging from -0.13-0.65. M7 generated the lowest NMB and RMSE and-likely due to cancelling effect of its overestimation for the summer and underestimation for other seasons (Fig. 3).

For NO, model correlations for EA1 ranged from 0.57-0.68, showing that all of models effectively reproduced spatial variability in NO for this subregion (Table 3). NMBs indicated that

10 underestimations by the models except in the case of M8, mostly occurred for the winter. This underestimation can be partly attributed to the coarse model horizontal resolution (45 km) used in MICS-Asia III, which hardly reproduced concentrations of short-lived species (e.g., NO). In contrast to most of the other models, M8 overestimated NO concentrations for all three subregions. It is noted that NO observations for EA3 were too low (<0.3 ppbv) to be discussed in this study.</p>

Table 4 shows statistics on the models' performance in measuring NO<sub>2</sub> levels. In general, most of the models performed better in representing NO<sub>2</sub> than O<sub>3</sub> and NO for EA1. NMBs ranged from -0.28-0.322, falling far below those measured for O<sub>3</sub> (0.48-1.25). Correlations of 0.54-0.66 were recorded, implying the models' reliable performance in reproducing spatial and monthly variability of NO<sub>2</sub> for EA1. Similar to those for O<sub>3</sub> and NO, correlation coefficients for NO<sub>2</sub> in EA2 remained low. Thus, a dedicated investigation of O<sub>3</sub>, NO and NO<sub>2</sub> levels in EA2 is urgently needed, but falls beyond the scope of this study. In EA3, correlation coefficients ranged from 0.543-0.72. NMBs and RMSEs except for those of M8 ranged from -0.42-0.23-0.46 and 0.901-1.79 ppby, respectively.

# 3.5 Vertical profiles of O<sub>3</sub>

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Fig. 5 shows the vertical profiles of observed and simulated O<sub>3</sub> levels for East Asia for the summer and winter. Ensemble means (Ense) show underestimations and overestimations of EA2 O<sub>3</sub> levels for the middle (500-800 hPa) and lower (below 900 hPa) troposphere, respectively. In the winter, underestimations extended to 200 hPa. Magnitudes of underestimations and overestimations reached 10-40 ppbv and 10-20 ppbv, respectively. For EA3, Ense reproduced the vertical structure of ozone for

both the summer and winter. An overestimation of less than 800 hPa, with a magnitude of 10-20 ppbv was observed for the summer.

High levels of intermodel variability in  $O_3$  exceeding 300 hPa is evident across all subregions, which is attributable to the varied top boundary conditions applied by the models. However, this

- 5 considerable variability was not transmitted to the middle troposphere (400-600 hPa), in which O<sub>3</sub> concentrations were consistent across the models. In the lower troposphere, a minor level of intermodel variability of below 900 hPa appeared in the winter in three sub-regions, and slowly decreased with height. Mean standard deviations (SD) of models of below 900 hPa were recorded as 7.6 ppbv, 6.9 ppbv and 6.0 ppbv for EA1, EA2 and EA3, respectively, covering 18.3%, 15.0% and 15.4% of mean
- O<sub>3</sub> concentrations. In 700-900 hPa, SD levels decreased to 5.4 ppbv, 4.4 ppbv and 4.8 ppbv for EA1,
   EA2 and EA3, 12.2%, 9.4% and 10.8% of mean O<sub>3</sub> concentrations, respectively.

In the lower troposphere, intermodel variability in the summer was generally higher than that in the winter. In polluted regions (EA1), SD levels reached 16.3 ppbv (20.8 % of mean concentrations) in the summer, greatly exceeding those in winter (6.2 ppbv, 15.2%). Various vertical structures of O<sub>3</sub>

- 15 were found below 700 hPa in summer. O<sub>3</sub> concentrations slowly increased with height in M8 and M11, but they mixed well in the PBL and decreased from 800 hPa to 700 hPa in the other models. Akimoto et al. (2019) found that the parameterization on downward O<sub>3</sub> transport from the upper boundary layer contributed considerably to discrepancies between M1, M6 and M11. In EA2, vertical structures of O<sub>3</sub> among models were found to be consistent, but concentrations differed more than those in EA1. SD
- 20 values covered 22% of mean concentrations.

#### 4. Multi-model ensemble O<sub>3</sub> and comparison with MICS-Asia II

### 4.1 Spatial distributions of single model and multi-model ensemble O3

Fig. 6 shows the spatial distributions of MICS-Asia III ensemble mean surface O<sub>3</sub> values (Ense) and the coefficient of variation (CV). The CV is defined as the standard deviation of the modeled O<sub>3</sub>
divided by the average. The larger the CV value, the lower the degree of consistency among the models. For the summer, Ense predicted an elevated O<sub>3</sub> concentration belt in the middle-latitudes (30°-45°N). A region of O<sub>3</sub> in excess of 60 ppbv stretched across the North China Plain and East China Sea,

far <u>exceeding</u> MICS-Asia II (45-50 ppbv) values for 2001 (Han et al., 2008). In other seasons, the  $O_3$  values distribution shows higher  $O_3$  over the ocean than in eastern China, reflecting  $O_3$  titration from high NO<sub>x</sub> emissions. Due to the stratospheric injection, surface  $O_3$  over the Tibet Plateau remained at high levels throughout the year, ranging from 50 to 65 ppbv. The seasonal cycle of surface  $O_3$  levels determined from Ense via MICS-Asia III agrees<u>d</u> with that observed from MICS-Asia II, but  $O_3$  levels in polluted regions were higher (Han et al., 2008).

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The CV ranged from 0.1-0.6 in East Asia. The highest\_values were found in EA1 in the winter. These high values in the low-latitude western Pacific (10°S-15°N) and Indian Oceans were likely caused by the treatment of lateral boundaries in the models. For MICS-Asia III, M7, M8 and M9

10 employed the default model configurations, and the others employed outputs of the GEOS-Chem/CHASER/MOZART-GOCART global model. Compared to those of MICS-Asia II, CVs for the Asian continent except for the winter remained at similar levels in this study (0.1-0.3) (Carmichael et al., 2008).

Although all of the models similarly predicted the emergence of an elevated summer O3

- 15 concentration belt in the middle-latitudes (30°-45°N), the magnitudes of enhanced O<sub>3</sub> levels varied between the models (Fig. 7). M5 predicted the highest O<sub>3</sub> concentrations of 60-90 ppbv for the North China Plain (EA1) and for its outflow pathways including the Bohai Sea, East China Sea, Korea, Japan and the Sea of Japan (Locations are shown in Fig. S1 in the supplementary section) whereas M8 predicted the lowest levels of 35-50 ppbv. Overhangs of 30 ppbv contour lines extending into the northwestern Pacific among along the Asian continent outflow plume differed considerably between the models. A plume of 30 ppbv or higher O<sub>3</sub> levels was observed simulated from in M1-M6, M13 and M14, reaching further south and east of Japan (135°E, 20°N), than those of M8, M10 and M11 (120°E, 30°N). From MICS-Asia II and HTAP, differences in the frequency of marine air masses from the western Pacific Ocean were thought to be a possible cause of O<sub>3</sub> discrepancies observed over oceans
- 25 between the models due to different meteorological drivers (Han et al., 2008). For MICS-Asia III, winds fields reflected employed by the models were similar due to the use of the same or similar meteorological fields (Fig. S2 in the supplementary section). These inconsistencies between the models resulted from the combined effects of a series of factors, including the diversity of condensed gas-

chemical mechanism and heterogeneous chemistry. Li et al. (20165) found chemical production to be the dominated controlling factor of O<sub>3</sub> along outflow pathways near the North China Plain in the summer rather than lateral and top boundary conditions. Impacts of aerosols on ozone O<sub>3</sub> in these regions are-were frequently reported in Olson et al. (1997) and Li et al. (2018) to alter photolysis rates and heterogeneous chemistry patterns. Detailed comparisons of parameterizations of these processes in models are needed in future intermodel comparison projects focused on Asia.

In the winter, distribution patterns of O<sub>3</sub> were quite similar between the models with high concentrations observed over parts of western China, northeastern India and the western Pacific from the East China Sea to southern Japan (Fig. S3 in the supplementary section). In <u>the</u> spring and autumn

10 (Fig. S4 and Fig. S5 in the supplements), O<sub>3</sub> concentrations were generally higher than they were in the winter across the whole model domain due to the enhancement of solar radiation or stratosphere-troposphere exchanging fluxes of O<sub>3</sub>. All of the models exhibited an enhancement of O<sub>3</sub> levels over southern Tibet, northeastern India and the western Pacific, generally echoing patterns observed in the winter. Increases in of O<sub>3</sub> observed in further north of Japan are-were comparable with winter.

#### 15 4.2 Comparison with MICS-Asia II

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From MICS-Asia II, model evaluation on O<sub>3</sub> were conducted on sites in the western Pacific. Fig. 8 presents the simulated and observed surface O<sub>3</sub> levels at these monitoring sites derived from the phase II and III of the MICS-Asia project. Note that different models were employed in two phases. In general, most of the models captured distributions of O<sub>3</sub> at most sites in both MICS-Asia II and III. ENSE results are were consistent for March and December of 2001 and 2010. Underestimations of O<sub>3</sub> levels in March at Japan sites (site 4: <u>OkiSado-seki</u>, site 5: <u>HedoOki</u> and site 6: Banryu) in Phase II were largely remedied in Phase III. However, surface O<sub>3</sub> observed in western Japan (site 4: Oki, site 5: Hedo and site 6: Banryu) were severely overestimated in July 2010 by 10-30 ppbv. This overestimation was not been-found in Phase II, for which differences from observations were valued at approximately 5 ppbv. Rural sites in western Japan are located in the upwind regions of Japanese domestic emissions, and are subjected to the impacts of Asian continent outflows. Overestimated O<sub>3</sub> values for North China Plain (EA1) in Phase III contributed considerably to enhanced concentrations measured <u>simulated</u> for sites of western Japan in July 2010. This indicates that transboundary transport from the Asian

continent according to MCIS-Asia III was likely overestimated relative to that measured from MICS-Asia II.

#### 5. Discussions

In reference to MICS-Asia II, Han et al. (2008) hypothesized that variations in meteorological
fields, dry deposition, PBL, model treatments of chemistry and other physical processes had contributed to model biases in relation to observations and intermodel variability. Quantifying the contributions of these processes can help explain model biases through sensitivity simulations. However, this task comes with tremendous computational costs when applied to 14 models. The qualitative analysis of potential causes of these processed processes based on comparisons of models
and observations is essential to selecting sensitivity simulating scenarios for the next phase of MICS-Asia. In MICS-Asia III, common input data (emission and meteorology) bewere effectively used in this qualitative analysis based on model parameterizations. We evaluated the models on dry depositions, PBL and chemistry by collecting their observations (dry deposition velocity and PBLH). This work was not conducted under MICS-Asia II and is intended to help model developers improve model

15 performance for East Asia.

### 5.1 Dry depositions

Previous studies show that dry deposition processes serve as the key net sink of O<sub>3</sub>, accounting for roughly 25% of total removed from the troposphere (Lelieveld and Dentener, 2000). The uncertainty of dry deposition in CTMs is still high because many processes are heavily parameterized in models (Hardacre et al., 2015). In this study, the simulated dry deposition velocities of O<sub>3</sub> were compared. Simulated deposition velocities were calculated from Eq. (1):

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 $V_d = F/C \tag{1}$ 

where F and C represent the simulated dry deposition flux and surface  $O_3$  concentrations, respectively. We determined spatial mean dry deposition velocities from stations in each subregion.

Fig. 9 presents the simulated and observed monthly spatial mean dry deposition velocities of O<sub>3</sub>.
 For EA1, ensemble mean values overestimated observed dry deposition velocities of O<sub>3</sub> (v<sub>d</sub>) for

August-September, but still fell within the range of the observed standard deviation. This shows that other factors rather than dry deposition could play important roles in overestimations of August-September O<sub>3</sub> values in EA1. In October-November, simulated v<sub>d</sub> apparently underestimated observations by 30-50%. Among the models, the lower dry deposition velocities in May-July for M1,

- 5 M2, M4 and M6 than that of M11 partly explained higher May-July surface O<sub>3</sub> from those simulations than that in M11. However, M13 and M14 still produced high O<sub>3</sub> concentrations in May-September although their dry deposition velocities were similar to that of M11(Fig. 3). Notably, our observations were made on grassland, which covers ~20% of the land area in EA1. There are few v<sub>d</sub> observations on agriculture crops (50% of the land area) in <u>North</u> China. Hardacre et al. (2015) reported O<sub>3</sub> dry
- deposition measurements on crops in Europe and simulated O<sub>3</sub> dry deposition in 15 global models.
   Both observations and simulations showed that O<sub>3</sub> dry deposition velocities on agriculture crop class were quite similar to those of grassland, showing uncertainties related to be the representativeness of measurement sites used in this study did not affect our conclusions.

For EA2, similar features as those of EA1 were found. M1, M2, M4 and M6 were quite consistent

- 15 with each other, with a seasonal cycle and a spring minimum. M11, M12 and M14 show no obvious signs of seasonal variability with a magnitude of 0.1-0.2 cm/s. Seasonal patterns in M13 are considerably different from those of the other models, exhibiting a maximum in April-September with higher dry deposition velocities (0.5 cm/s). The performance of the models for dry deposition velocities was not always consistent with O<sub>3</sub> concentrations. For example, O<sub>3</sub> concentrations in M13 remained
- 20 high under higher dry deposition velocities.

In EA3, most stations were located in remote oceanic sites, and few dry deposition observations were made. Thus, we collected observations from other oceanic sites to evaluate model performance (Helmig et al., 2012). Ense values for  $v_d$  agree<u>d</u> reasonably well with observations (Fig. 9). Both observations and simulated  $v_d$  values show<u>ed</u> a July-September maximum with a magnitude of 0.02-

25 0.03 cm/s. Park et al. (2014) found surface O<sub>3</sub> levels in EA3 to be more sensitive to dry deposition parameterization schemes in CTMs. O<sub>3</sub> measured from oceans differed by 5-15 ppbv in East Asia due to the use of various dry deposition parameterization schemes. Thus, more observations are needed over oceans in EA3 to mitigate O<sub>3</sub> simulation uncertainties.

#### 5.2 Relationships between surface NO<sub>x</sub> and O<sub>3</sub>

In general, surface  $O_3$  mainly comes from photochemistry processes involving  $NO_x$  and VOCs in polluted regions. Examining  $O_3$ - $NO_x$  relationships is effectual to investigating sources of intermodel variability and model errors concerning  $O_3$  chemistry in East Asia. Fig. 10 presents  $O_3$  concentrations as a function of  $NO_x$  in May-September based on the monthly daytime (8:00-20:00) mean observed and

simulated results for the stations shown in Fig. 1.

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For EA1 (North China Plain), observations clearly show that  $O_3$  concentrations decreased with an increase in NO<sub>x</sub> concentration.  $O_3$  concentrations mostly remained at high levels (40-60 ppbv) when NO<sub>x</sub> was less than 20 ppbv. The slope and intercept of the regression line between observed  $O_3$  and

- 10 NO<sub>x</sub> were measured as -0.77 ppbv/ppbv and 59.5 ppbv, respectively. Among the models, M11 results are-were in relative agreement with observations. The slope and intercept (-1.01 ppbv/ppbv, 63.23 ppbv) reflected the observations. The other models showed a higher degree of model bias and intermodel variability in relationships between O<sub>3</sub> and NO<sub>x</sub>. Their slopes mostly ranged from -1.25 ppbv/ppbv to -2.13 ppbv/ppbv, amounting to 1.3-2.8 times of observed slope. Their intercepts were
- 15 74.9 -121.2 ppbv, far exceeding observations (59.5 ppbv). Akimoto et al. (2019) calculated the net photochemical production of M1, M6 and M11 and found that weak net chemical production in M11 were mostly responsible for low O<sub>3</sub> rather than those in M1 and M6. This finding is consistent with the low slope in M11. To reduce the impact of O<sub>3</sub> buildup and transport by consuming NO<sub>x</sub>, relationships between O<sub>x</sub> (NO<sub>2</sub>+O<sub>3</sub>) with NO<sub>x</sub> was compared (Fig. S7 in the supplementary section). Observed O<sub>x</sub>
- 20 increases with the increase of NO<sub>x</sub> levels, with coefficient of determination (R<sup>2</sup>) of 0.61. Most of the models (except for M8, M11 and M13) failed to reproduced observed positive correlations between O<sub>x</sub> and NO<sub>x</sub>, and their R<sup>2</sup> only ranged from 0.01-0.08. The slope, intercept and R<sup>2</sup> of M8 and M11 arewere relative agreement with observations.

For EA2, all models reproduced observed key patterns in which Ox positively correlated with

NO<sub>x</sub>. For O<sub>3</sub>-NO<sub>x</sub> relationships, M1, M2, M4 and M6 reproduced observed O<sub>3</sub> levels under low NO<sub>x</sub> conditions(conditions (< 30 ppbv) but failed to capture low O<sub>3</sub> under high NO<sub>x</sub> conditions (30–40 ppbv), accounting for overestimations of these models for O<sub>3</sub> in May-September. By contrast, M8 and M11 produced excessively high NO<sub>x</sub> values, resulting in their underestimations of O<sub>3</sub> values. For M13

and M14,  $O_3$  concentrations were nearly constant in all levels of  $NO_x$ .  $O_3$  was positively correlated with  $NO_x$  in M9 and M10, which stands in contrast to observations. This finding suggests that more attentions are is needed using M9, M10, M13 and M14.

Stations in EA3 are mostly located over clean oceans or islands. NO<sub>x</sub> concentrations were less than 3 ppbv, showing that local chemistry was not a key factor shaping O<sub>3</sub> formation. Thus, we did not examine the simulated O<sub>3</sub>-NO<sub>x</sub> relationship further.

#### 5.3 Other factors

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Previous studies show that  $O_3$  precursors are mostly constrained within the boundary layer (Quan et al., 2013). PBLH model evaluation is essential for the interpretation of model biases with

- 10 observations. Unfortunately, this evaluation was not applied in MICS-Asia II. In 2016, Guo et al. (19962016) calculated the PBLH using the bulk Richardson number (Ri) method from the radiosonde network of the L-band sounding system of the China Meteorological Administration (Vogelezang and Holtslag, 1996). The system provides fine-resolution profiles of temperature, pressure relative humidity, wind speed and direction. In MICS-Asia III, all selected models exhibited the spring-
- 15 maximum and winter-minimum season cycle for EA1 (Fig. S6 in the supplementary section), capturing the main climatological pattern of PBLH observations (Guo et al., 2016). The Ense on PBLH only overestimated radiosonde measurements by 100-200 m (~10-15%) and likely due to sampling bias between the models and measurements. The simulation was recorded as the mean value of 12 hours (08:00-20:00), while the average of the measurements was calculated based on a 3-hour period (08:00,
- 20 14:00 and 20:00). For EA2, the observed PBLH did not vary as much as that for EA1, and differences between seasons ranged within 100 m. This pattern was captured by the models. As was observed from EA1, the simulated PBLH for EA2 exceeded the measurements by100-200 m. Few measurements of remote oceanic sites in East Asia were collected. Thus, we compared simulations with European Centre for Medium-Range Weather Forecasts Reanalysis Data (von Engeln et al., 2013). Both showed a
- 25 winter-maximum pattern for PBLH.

#### 6. Summary

Under MICS-Asia III framework, the evaluation and intercomparison of 13 CTMs was-were conducted using a wide variety of observations covering two Chinese industrialized regions and the western Pacific, using long-termcontinuous simulations for 2010 with a focus on O<sub>3</sub>, NO and NO<sub>2</sub>. In

5 particular, surface O<sub>3</sub> levels in China, which were neglected in previous model-intercomparison projects, were evaluated. Considerable levels of intermodel variability in O<sub>3</sub> were observed across all subregions of East<u>Asia</u>, with model concentrations varying by factors of 2 to 3 between different models.

A model ensemble was produced and evaluated. In general, the model ensemble captured key patterns of monthly and diurnal O<sub>3</sub>, NO and NO<sub>2</sub> in the North China Plain and western Pacific Rim. It failed to capture the observed seasonal cycle of O<sub>3</sub> for the Pearl River Delta of China. For the North China Plain and western Pacific Rim, the model ensemble severely overestimated surface O<sub>3</sub> levels for May-September by 10-30 ppbv. This overestimation systematically appeared in both daytime and nighttime. Similarly, the model ensemble tended to overestimate spring daytime and nighttime O<sub>3</sub>

- 15 concentrations for the Peral River Delta. Compared to MICS-Asia II, MICS-Asia III was less prone to underestimating surface O<sub>3</sub> in March for Japanese sites. However, it predicted excessively high surface O<sub>3</sub> concentrations for western Japan in July, which was not the case for MICS-Asia II. In term of O<sub>3</sub> soundings, the ensemble model used in this study reproduced the vertical structure in the western Pacific, but overestimated O<sub>3</sub> below 800 hPa in the summer. For the industrialized Pearl River Delta,
- 20 the ensemble average presented an overestimation of O<sub>3</sub> levels for the lower troposphere and underestimations in the middle troposphere. We find that the ensemble average of 13 models for O<sub>3</sub> does not always perform better than individual models for East Asia in contrast to their performance for Europe. This suggest that the spread of ensemble-model values does not represent all uncertainties in O<sub>3</sub> levels or that most MICS-Asia III models missed key processes. In contrast to performance levels

for O<sub>3</sub>, <u>ENSE Ense performed</u> –better than individual models for NO<sub>2</sub> in East Asia.

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MICS-Asia II outlines potential causes of variability among models. Quantifying the contributions of these processes to O<sub>3</sub> concentrations serves as an effective way to explain model biases through sensitivity simulations. However, this would incur tremendous computational costs when

applied to 14 models. In this study, we conducted a qualitative analysis of potential causes by comparing models and observations for these processes to identify sensitivity simulating scenarios for the next phase of MICS-Asia. Our comparisons show that the ensemble model overestimated observed dry deposition velocities of O<sub>3</sub> for August-September in North China Plain, showing that other factors

- 5 rather than dry deposition may contribute to the overestimation of simulated O<sub>3</sub> concentrations in the summer. For the western Pacific, simulated v<sub>d</sub> values agreed with observations reasonably well. Photochemical treatment in models may contribute to O<sub>3</sub> overestimations in North China Plain. The studied models captured major <u>climatological</u> pattern of\_<u>climatological</u> pattern of PBLH observations for three subregions of East Asia. More evaluations of turbulent kinetic energy in PBL are needed to
- 10 assess vertical mixing in future studies.

# Author contribution:

JL, ZW and GC conducted the study design. JL, TN, BG, KY, JF, XW, QF, SI, HL, CK, CL, MZ, ZT, MK, HL contributed to modeling data. ML, JW, JK and QW provided the emission data. LK helped with data processing. HA, GC and ZW were involved in the scientific interpretation and

15 discussion. JL

prepared the manuscript with contributions from all co-authors.

# **Competing interests**:

The authors declare that they have no conflict of interest.

# Acknowledgements:

This work was supported by the Natural Science Foundation of China (41620104008;
 41571130034; 91544227; 91744203), and National Key R&D Program of China (2017YFC0212402).
 This work was partly supported by the Environment Research and Technology Development Fund (S-12) of the Environmental Restoration and Conservation Agency of Japan and the Ministry of Environment, Japan. We thank the Pearl River Delta Regional Air Quality Monitoring Network for

25 observations in Pearl River Delta. Dr. Kengo Sudo from Nagoya university and Prof. Rokjin J. Park provided us CHASER and GEOS-Chem outputs for boundary conditions. This manuscript was edited by Wallace Academic Editing.

# **Appendix A. Statistical Measures**

Defining  $y_{ij}$  and  $Obs_{ij}$  modeled and observed the i<sup>th</sup> monthly concentrations of air pollutants at the j<sup>th</sup> station, having mean value  $\bar{y}$  and  $\overline{obs}$ . m and n represent the numbers of stations and months.

Correlation coefficient (R)

$$R = \frac{\sum_{j=1}^{m} \sum_{i=1}^{n} (y_{ij} - \bar{y})(obs_{ij} - \overline{obs})}{\sum_{j=1}^{m} \sum_{i=1}^{n} (y_{ij} - \bar{y})^2 \sum_{j=1}^{m} \sum_{i=1}^{n} ((obs_{ij} - \overline{obs})^2)}$$
(A1)

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Root mean square error (RMSE):

$$\text{RMSE} = \sqrt{\frac{\sum_{j=1}^{m} \sum_{i}^{n} (y_{ij} - Obs_{ij})^2}{n}}$$
(A2)

Normalized Mean Bias (NMB)

$$NMB = \frac{\sum_{j=1}^{m} \sum_{i=1}^{n} (y_{ij} - Obs_{ij})}{n \times \bar{y} \times \overline{Obs}}$$
(A3)

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## **Table and Figure captions:**

Table.1 Basic structures, schemes and relevant parameters of the fourteen participating models

Table. 2 Statistical analysis for surface O<sub>3</sub> in three subregions over East Asia (R: correlation coefficient;
10 NMB: Normalized Mean Bias; RMSE: Root Mean Square Error)

Table. 3 Statistical analysis for surface NO in three subregions over East Asia (R: correlation coefficient; NMB: Normalized Mean Bias; RMSE: Root Mean Square Error)

Table. 4 Statistical analysis for surface NO<sub>2</sub> in three subregions over East Asia (R: correlation coefficient; NMB: Normalized Mean Bias; RMSE: Root Mean Square Error)

- 15 Fig. 1 Model domain of models for except M13 and M14 with the locations of three subregions marked in this study. Also shown are the locations of surface monitoring stations used in this study. The meteorological model used to provide meteorological fields for most models also uses this domain. Note that the domains of M13 and M14 are shown in Fig.10.
- Fig. 2 Box-plots of observed and simulated annual NO<sub>2</sub> (left column), NO (middle column) and O<sub>3</sub> (right column) frequency distribution determined from 13 models and averaged for stations in EA1, EA2 and EA3 in time for 2010. n denotes the numbers of stations. The rectangle represents the interquantile range (25<sup>th</sup> to 75<sup>th</sup> percentiles). The small star identifies the mean, the continuous horizontal line within the rectangle identifies the median, and whiskers extend between minimum and maximum values.
- Fig. 3 Time series of monthly NO<sub>2</sub>, NO and O<sub>3</sub> levels simulated by all models and their ensembles (Ense)
  in ppbv, averaged over all observed stations across three subregions of East Asia (EA1: top row, EA2: middle row, EA3: bottom row). Observations are denoted by the black line. n represents the number of stations. <u>Gray lines represent NO<sub>2</sub>, NO and O<sub>3</sub> levels simulated by models except M1, M2, M4, M6, M11, M12, M13 and M14.
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Fig. 4 Seasonal mean diurnal cycle of surface O<sub>3</sub>, in ppbv, as a function of hours, for all models and their
ensembles, averaged across all observed stations in three subregions of East Asia (EA1: top row, EA2: middle row, EA3: bottom row). Observations are denoted by the black line. n represents the number of

stations. Gray lines represent O<sub>3</sub> levels simulated by models except M1, M2, M4, M11, M12 and M14. Spring, summer, autumn and winter were defined as time periods of March-April-May, June-July-August, September-October-November, and December-January-February, respectively.

Fig. 5 Simulated O<sub>3</sub> profiles for the summersummer (June-July-August) and winter (December-January<u>February</u>) of 2010, averaged over all observed stations across three subregions of East Asia (EA1: left column, EA2: middle column, EA3: bottom column). Ozonesonde data for 2010 were taken from World Ozone and Ultraviolet Radiation Data Centre (WOUDC) database

Fig. 6 Ensemble mean seasonal surface O<sub>3</sub> concentrations and CV values for different seasons. CV is defined as the standard deviation of the modeled fields divided by the average for different seasons

Fig. 7 Surface O<sub>3</sub> spatial distribution derived from 13 models for summer 2010 (unit: ppbv).

Fig. 8 Modeled and observed monthly mean concentrations of O<sub>3</sub> for EANET sites in the phase II (left panel) and III (right panel) of the MICS-ASIA project. The solid line represents the ensemble mean. Note that MCIS-ASIA II and III data refer to March, July and December of 2001 and 2010, respectively. IDs

of the monitoring sites denote the following: 1: Rishiri (45.12°N, 141.23°E), 2:Ogasawara (27.83°N, 142.22°E), 3: Sado-seki (38.23°N, 138.4°E), 4: Oki (36.28°N, 133.18°E), 5: Hedo (26.85°N,128.25°E),
6: Banryu (34.67°N,131.80°E)

Fig. 9 Simulated and observed monthly  $O_3$  dry deposition velocities ( $V_d$ ) for M1, M2, M4, M6, M11, M12, M13 and M14 for three subregions of East Asia (EA1: top row, EA3: middle row, EA3: bottom

- 20 row). TEX, STR, GGSEX and AMMA denote observations for TexAQS06 (7 July–12 September 2006; north-western Gulf of Mexico), STRATUS06 (9–27 October 2006; the persistent stratus cloud region off the coast of Chile in the eastern Pacific Ocean), GasEx08 (29 February– 11 April 2008; the Southern Ocean), and AMMA08 (27 April–18 May 2008; the southern and northern Atlantic Ocean). Observational data were taken from Sorimachi et al. (2003), Pan et al. (2010), and Helmig et al. (2012).
- Fig. 10 Scatter plots for monthly daytime (08:00-20:00) surface NO<sub>x</sub> and O<sub>3</sub> for each station in EA1 (red), EA2 (green)and EA3 (blue) in May-October, for observations (obs) and models. Also shown are the linear regression equations for NO<sub>x</sub> and O<sub>3</sub> in EA1 (red) and EA2 (green).

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Models	M1	M2	M3	M4	M5	M6	M7	M8	M9	M10	M11	M12	M13	M14
Domain	Ref <sup>a</sup>	Global	10 °N -50°N;											
														80 °E -135 °E
Horizontal	45 km	0.5 ° ×0.667°	45 km											
resolution														
Vertical	$40\sigma_p$ levels	$60\sigma_p$ levels	$20\sigma_z$ levels	$40\sigma_p$ levels	$47\sigma_p$ levels	$15\sigma_z$ levels								
resolution														
Depth of first	58 m	29 m	58 m	16 m	44 m	48 m	27 m		100 m					
layer														
Meteorology	Standard <sup>b</sup>	WRF/NCEP <sup>b</sup>	WRF/NCEP <sup>b</sup>	WRF/NCEP <sup>b</sup>	WRF/	Standard <sup>b</sup>	Standard <sup>b</sup>	GEOS-5	RAMS/NCEP <sup>b</sup>					
										MERRA2 <sup>b</sup>				
Advection	Yamo	Yamo	Yamo	PPM(Colle	PPM	Yamo	5 <sup>th</sup> order	5 <sup>th</sup> order	5 <sup>th</sup> order	5 <sup>th</sup> order	Walcek	Walcek and	PPM	PPM
	(Yamartino,			lla and			monotonic	monotonic	monotonic	monotonic	and	Aleksic		
	1993)			Woodward							Aleksic	(1998)		
				1984)							(1998)			
Vertical	ACM2	ACM2	ACM2	ACM2	ACM2	ACM2	3 <sup>th</sup> order	3 <sup>th</sup> order	YSU	YSU	K-theory	FTCS	Lin and	ACM2
diffusion	(Pleim,2007)						Monotonic	Monotonic				(Forward in	McElroy,	
												Time, Center	(2010)	
												in Space)		
Dry	Wesely	Wesely	Wesely	M3DRY	M3DRY	M3DRY	Wesely	Wesely	Wesely	Wesely	Wesely	Wesely(1989	Wesely	Wesely (1989
deposition	(1989)	(1989)	(1989)	(Pleim et			(1989)	(1989)	(1989)	(1989)	(1989)	)and Zhang	(1989)	
				al., 2001)								et al. (2003)		

Table1 Basic structures, schemes and relevant parameters of the fourteen participating models

Wet	Henry's Law	Henry's	Henry's	Henry's	Henry's Law	ACM	Henry's Law	AQCHEM	Easter et al.,	Grell	Henry's	Henry's Law	Henry's Law	Henry's Law
deposition		Law	Law	Law					(2004)		Law			
Gas	SAPRC99(C	SAPRC99	CBM05(Y	SAPRC99	SAPRC99	SAPRC99	RACM-	RACM	RADM2	RADM2	CBMZ	SAPRC99(C	NOx-Ox-HC	SAPRC99
chemistry	arter,2 000)		arwood et				ESRL with	(Goliff et	(Stockwell et		(Zaveri et	arter,2000)	chemistry	
			al.,2005)				KPP	al., 2013)	al., 1990)		al.,1999)		mechanism	
Aqueous	ACM-ae6	ACM-ae6	ACM-ae5	ACM-ae5	ACM-ae5	ACM-ae5	CMAQ	AQCHEM	Walcek and	None	RADM2	Walcek and	-	ACM
chemistry							simplified		Taylor		(Stockwell	Teylor		
							Aqueous		(1986)		et al.,	(1986)		
							chemistry				1990)	Carlton et al.		
												(2007)		
Inorganic	AER06(Bink	AER06	AER05	AER05	AER05	AER05	MADE	MADE	MADE	GOCART	ISORROP	Kajino et al.	ISORROPIAv1.	ISORROPIAv1.7
mechanism	owski and						(Ackermann				IAv1.7(Ne	(2012)	7	
	Roselle,						et al., 1998)				nes et			
	2003)										al.,1998)			
Boundary	GEOS-Chem	Gipson	GEOS-	CHASER	CHASER	CHASER	Liu et al.	CHASER	GEOS-Chem	MOZART	CHASER	CHASER	/	GEOS-Chem
conditions	global model	(1999)	Chem	global	global model	global	(1996)	global model	global model	+	global	global model		global model
	(Martin et		global	model		model				GOCART	model			
	al.,2002)		model	(Sudo et						global				
				al., 2002a,						models <sup>c</sup>				
				2002b)										
Two-way	Off-line	Off-line	Off-line	Off-line	Off-line	Off-line	On-line	On-line	On-line	Off-line	Off-line	On-line	Off-line	Off-line
feedback														

<sup>a</sup> Ref represent the referenced domain by MICS-ASIA III project.

<sup>b</sup>Standard represents the reference meteorological field provided by MICS-ASIAIII project; WRF/NCEP and WRF/MERRA represents the meteorological field of the participating model itself, which was run by WRF driven by the NCEP and Modern Era

Retrospective-analysis for Research and Applications (MERRA) reanalysis dataset.

Boundary conditions of M10 are from MOZART and GOCART (Chin et al., 2002; Horowitz et al., 2003), which provided results for gaseous pollutants and aerosols, respectively.

Models	Region	R	NMB	RMSE	Region	R	NMB	RMSE	Region	R	NMB	RMSE
M1		0.89	0.52	19.79		0.48	0.31	14.41		0.57	0.28	15.49
M2		0.90	0.64	18.13		0.10	0.35	15.06		0.66	0.24	13.83
M4		0.87	0.44	18.78		0.41	0.36	14.15		0.01	0.05	17.57
M5		0.87	0.42	19.00		0.30	0.14	13.38		0.34	0.31	19.28
M6		0.90	0.88	25.41		0.15	0.44	17.41		0.52	0.31	16.52
M7	EA1 (n=19) <sup>a</sup>	0.84	0.25	10.03	EA2 (n=13)	0.29	-0.08	11.11	EA3 (n=8)	0.60	0.02	10.97
M8		0.78	-0.47	13.52		0.20	-0.59	19.54		0.55	-0.27	15.32
M9		0.85	0.59	14.84		0.63	0.48	15.69		0.26	-0.09	13.27
M10		0.82	1.24	32.70		0.51	0.72	21.71		0.52	0.11	12.68
M11		0.81	0.09	9.46		0.34	-0.25	13.40		0.65	0.15	12.09
M12		0.89	0.55	18.53		0.36	0.30	13.31		0.57	0.11	11.81

Table 2 Statistical analysis for surface O<sub>3</sub> in three subregions over East Asia (R: correlation coefficient; NMB: Normalized Mean Bias; RMSE: Root Mean Square Error, unit is ppbv)

M13	0.86	0.95	22.69	0.25	0.50	17.04	0.63	0.09	11.04
M14	0.86	0.75	23.33	0.12	0.40	17.01	-0.13	-0.30	20.03
Ensemble Mean	0.89	0.53	15.92	0.38	0.23	11.76	0.52	0.08	11.93
Ensemble Media	0.89	0.56	17.86	0.37	0.31	13.29	0.54	0.11	12.06

a: n represents the numbers of observation stations

M1 $0.58$ $-0.35$ $20.68$ $0.22$ $-0.81$ $15.16$ $0.03$ $-0.35$ M2 $0.57$ $-0.14$ $23.73$ $0.14$ $-0.73$ $15.21$ $0.06$ $-0.27$ M4 $0.60$ $-0.61$ $22.29$ $0.18$ $-0.87$ $15.72$ $0.00$ $-0.39$ M5 $0.57$ $-0.07$ $20.34$ $0.24$ $-0.29$ $13.80$ $0.02$ $0.08$ M6EA1 $0.60$ $-0.71$ $23.36$ EA2 $0.11$ $-0.89$ $15.94$ $EA3$ (n=8)M7 $(n=19)$ $0.63$ $-0.75$ $24.91$ $(n=13)$ $0.04$ $-0.78$ $15.32$ $0.27$ $-0.40$ M8 $0.65$ $0.91$ $26.89$ $0.29$ $1.14$ $25.06$ $0.24$ $3.53$ M9 $0.58$ $-0.82$ $27.73$ $0.32$ $-0.93$ $16.72$ $0.22$ $-0.54$	Aodels	Region	R	NMB	RMSE	Region	R	NMB	RMSE	Region	R	NMB	RMSE
M4       0.60       -0.61       22.29       0.18       -0.87       15.72       0.00       -0.39         M5       0.57       -0.07       20.34       0.24       -0.29       13.80       0.02       0.08         M6       EA1 (n=19)       0.60       -0.71       23.36       EA2 (n=13)       0.11       -0.89       15.94       EA3 (n=8)       0.15       -0.70         M7       0.63       -0.75       24.91       0.04       -0.78       15.32       EA3 (n=8)       0.24       -0.29       1.14       25.06       0.24       3.53	<i>A</i> 1		0.58	-0.35	20.68		0.22	-0.81	15.16		0.03	-0.35	0.23
M5 $0.57$ $-0.07$ $20.34$ $0.24$ $-0.29$ $13.80$ $0.02$ $0.08$ M6EA1 $0.60$ $-0.71$ $23.36$ EA2 $0.11$ $-0.89$ $15.94$ $-0.79$ $0.15$ $-0.70$ M7 $(n=19)$ $0.63$ $-0.75$ $24.91$ $(n=13)$ $0.04$ $-0.78$ $15.32$ $0.27$ $-0.40$ M8 $0.65$ $0.91$ $26.89$ $0.29$ $1.14$ $25.06$ $0.24$ $3.53$	<i>A</i> 2		0.57	-0.14	23.73		0.14	-0.73	15.21		0.06	-0.27	0.19
M6EA1 (n=19)0.60-0.7123.36EA2 (n=13)0.11-0.8915.94 (n=13)0.15-0.70 (n=13)M7 $(n=19)$ 0.63-0.7524.91 $(n=13)$ 0.04-0.7815.320.27-0.40M80.650.9126.890.291.1425.060.243.53	<i>/</i> 14		0.60	-0.61	22.29		0.18	-0.87	15.72		0.00	-0.39	0.20
EA1EA2EA3 (n=8)M7 $(n=19)$ $0.63$ $-0.75$ $24.91$ $(n=13)$ $0.04$ $-0.78$ $15.32$ $0.27$ $-0.40$ M8 $0.65$ $0.91$ $26.89$ $0.29$ $1.14$ $25.06$ $0.24$ $3.53$	<i>A</i> 5		0.57	-0.07	20.34		0.24	-0.29	13.80		0.02	0.08	0.35
M7       (n=19)       0.63       -0.75       24.91       (n=13)       0.04       -0.78       15.32       0.27       -0.40         M8       0.65       0.91       26.89       0.29       1.14       25.06       0.24       3.53	<i>A</i> 6	EA1	0.60	-0.71	23.36	EA2	0.11	-0.89	15.94		0.15	-0.70	0.16
	<i>A</i> 7	(n=19)	0.63	-0.75	24.91	(n=13)	0.04	-0.78	15.32	EA3 (n=8)	0.27	-0.40	0.15
M9 0.58 -0.82 27.73 0.32 -0.93 16.72 0.22 -0.54	/18		0.65	0.91	26.89		0.29	1.14	25.06		0.24	3.53	0.94
	/19		0.58	-0.82	27.73		0.32	-0.93	16.72		0.22	-0.54	0.14
M10 0.63 -0.90 27.97 0.27 -0.94 16.30 0.39 -0.51	<i>A</i> 10		0.63	-0.90	27.97		0.27	-0.94	16.30		0.39	-0.51	0.14
M11 0.61 -0.34 19.92 0.04 -0.05 14.86 0.41 0.09	<b>/</b> 111		0.61	-0.34	19.92		0.04	-0.05	14.86		0.41	0.09	0.14

Table 3 Statistical analysis for surface NO in three subregions over East Asia (R: correlation coefficient; NMB: Normalized Mean Bias; RMSE: Root Mean Square Error, unit is ppbv)

M12	0.62	-0.55	21.19	0.13	-0.85	15.64	0.17	-0.48	0.16
M13	-	-	-	-	-	-	-	-	-
M14	0.68	-0.66	22.74	0.01	-0.66	14.77	0.24	-0.50	0.15
Ensemble Mean	0.63	-0.42	20.12	0.21	-0.55	13.58	0.20	-0.03	0.19
Ensemble Media	0.62	-0.58	21.66	0.17	-0.83	15.40	0.17	-0.45	0.16

a: n represents the numbers of observation stations

Models	Region	R	NMB	RMSE	Region	R	NMB	RMSE	Region	R	NMB	RMSE
M1		0.59	-0.18	11.08		0.33	-0.30	12.92		0.54	0.27	1.51
M2		0.64	-0.25	11.30		0.25	-0.43	14.85		0.43	-0.07	1.13
M4		0.65	-0.28	11.62		0.26	-0.32	13.79		0.56	-0.07	1.04
M5		0.57	0.08	10.86		0.30	0.09	12.91		0.60	0.46	1.79
M6	EA1	0.65	-0.22	11.04	EA2	0.23	-0.30	13.86	EA3 (n=8)	0.56	-0.23	0.90
M7	(n=19)	0.59	-0.22	11.42	(n=13)	0.20	-0.25	13.24		0.65	0.19	1.42
M8		0.43	<u>0.</u> 14 <del>.32</del>	11.90		0.43	0.15	10.97		0.72	2.38	4.46
M9		0.60	<u>0.</u> 32 <del>.30</del>	18.80		0.51	-0.37	12.66		0.49	0.05	1.66
M10	0.6	0.61	<u>0.11</u> - <del>10.61</del>	10.65		0.15	-0.08	12.81		0.63	0.06	1.33

Table 4 Statistical analysis for surface NO<sub>2</sub> in three subregions over East Asia (R: correlation coefficient; NMB: Normalized Mean Bias; RMSE: Root Mean Square Error, unit is ppbv)

M11	0.54	0.00	10.82	0.24	0.13	13.56	0.69	0.36	1.58
M12	0.63	-0.16	10.76	0.25	-0.24	13.78	0.61	-0.05	0.91
M13	-	-	-	-	-	-	-	-	-
M14	0.66	-0.12	10.00	0.08	-0.22	14.50	0.60	0.42	0.91
Ensemble Mean	0.65	-0.09	9.89	0.29	-0.18	12.16	0.64	0.25	1.33
Ensemble Media	0.65	-0.13	10.07	0.27	-0.23	12.85	0.59	0.06	1.23

a: n represents the numbers of observation stations

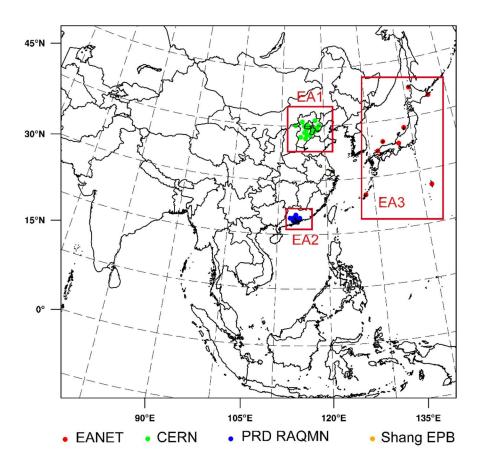


Fig.1 Li et al., 2018

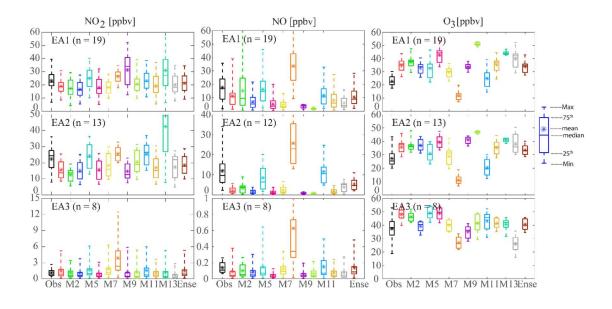


Fig.2 Li et al., 2018

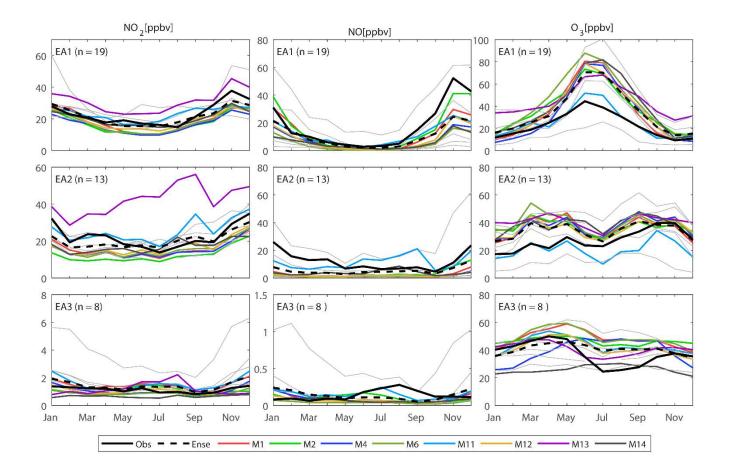


Fig.3 Li et al., 2018

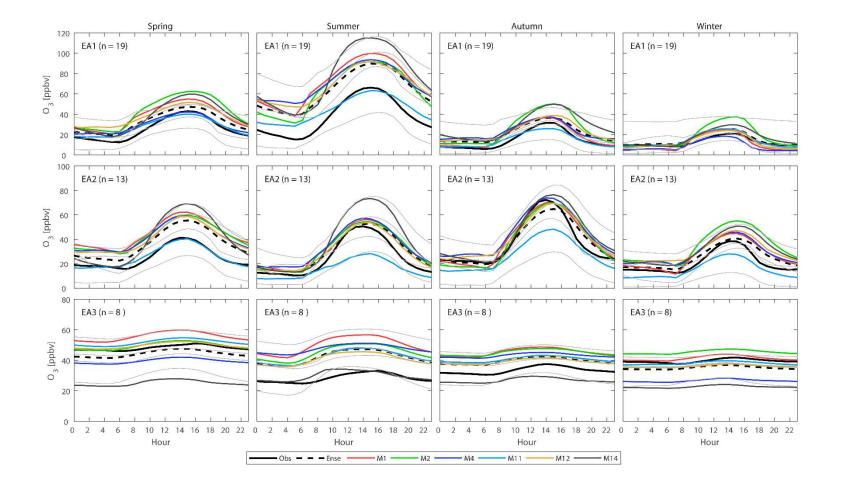
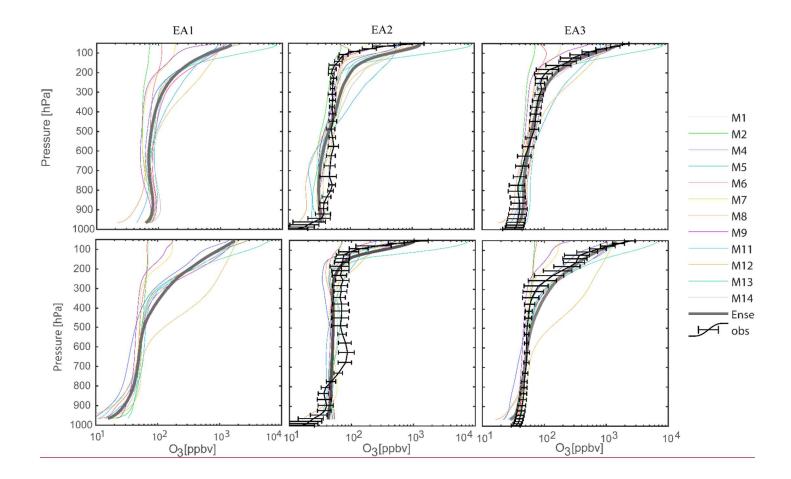


Fig.4 Li et al., 2018



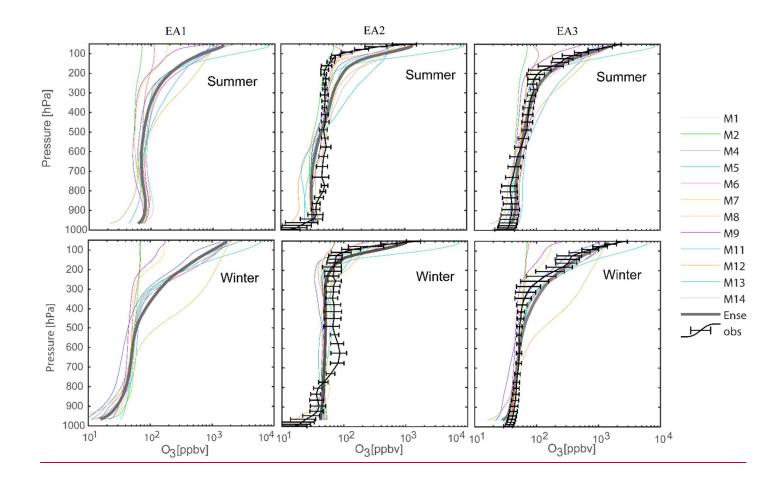


Fig.5 Li et al., 2018

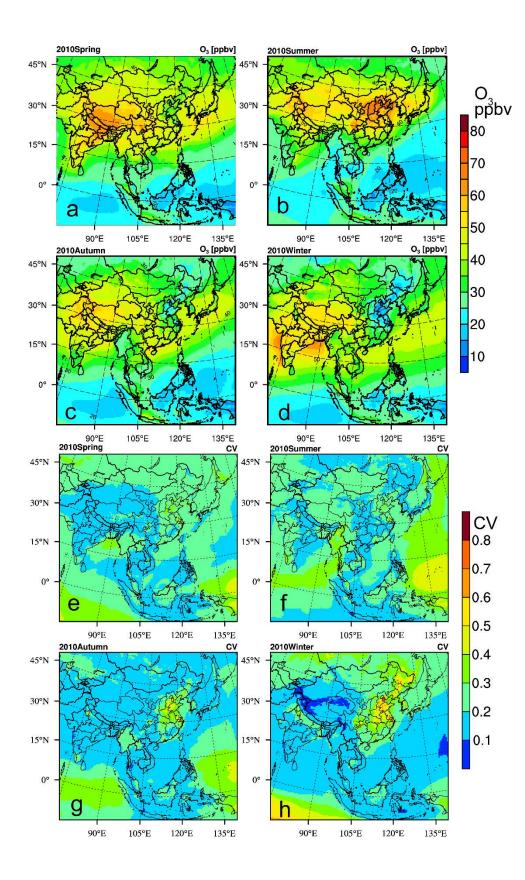


Fig.6 Li et al., 2018

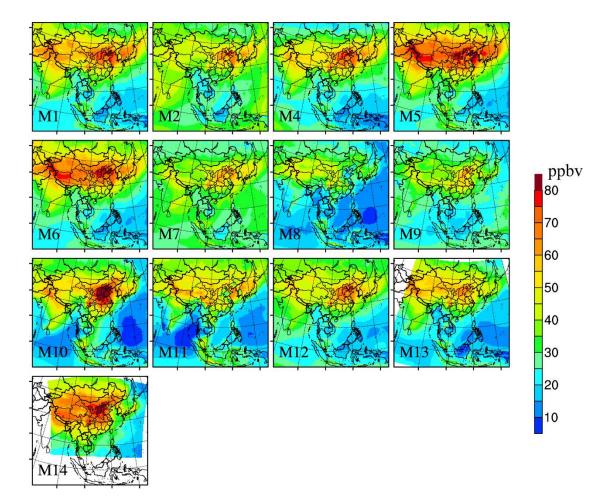


Fig.7 Li et al., 2018

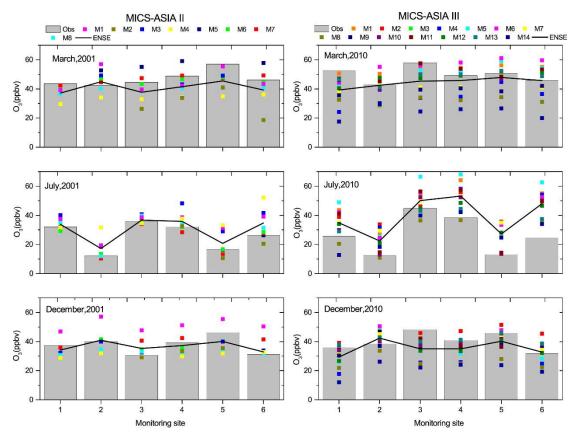


Fig.8 Li et al., 2018

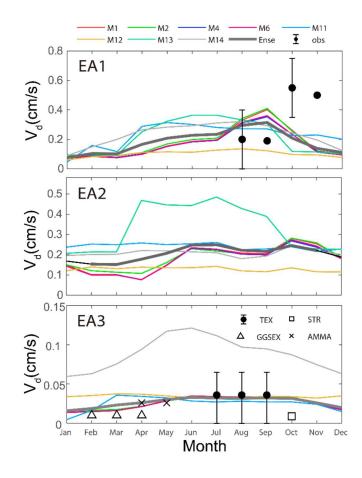


Fig.9 Li et al., 2018

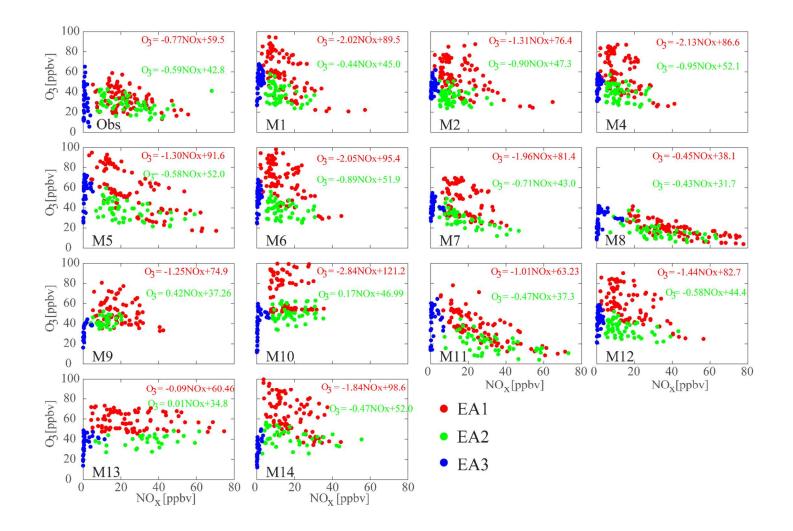


Fig.10 Li et al., 2018