We thank the reviewer for his/her constructive comments. Response to the Specific comments.

Comment 1: What does the vertical axis on the right of Fig. R1 mean?

If it represents a 3 km mesh emission, contrary to what is stated in the response document or in the revised manuscript, it is likely that the observation points receive local emissions.

Reply: Before the intercomparison between emission rates in 45 km and 3 km emission inventory, we provide the equation how to calculate the emission rates ($\mu g/m^2/s$) in 45 km resolution (E_{45km}) from those in 3 km resolution (E_{3km}).

$$E_{45km} = \sum_{i=1}^{m=15} \sum_{j=1}^{n=15} (E_{3km \, ij} / (m \times n)) \quad (1)$$

As shown in Eq (1), some $E_{3km\,ij}$ were higher than E_{45km} . Others could be less than E_{45km} . In general, there could be local emissions in ith and jth grid when E_{3km} in this grid is much higher than E_{45km} . When E_{3km} is much less than E_{45km} , the ith and jth grid represented a regional background condition. If E_{3km} is close to E_{45km} , the emission rates in the ith and jth grid represented the averaged conditions of 45 km×45 km areas. So, the comparison of E_{3km} with E_{45km} is a good way to examine if this 3km grid receives local emissions.

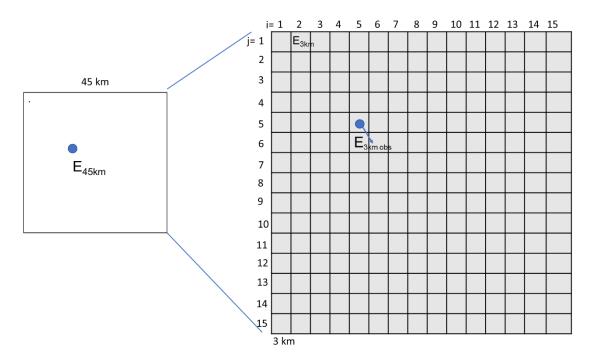


Fig. R1 The grids of observations in 45 (left) and 3 km (right) emissions inventory. The solid cycle represents the location of observation site.

In this study, we calculated the NO emission rates of observation sites in 45 km grid (E_{45km}). Emission rates (E_{3km}) at each 3 km grid (i=1,15; j=1,15) within this 45km grid was compared with E_{45km} . Meanwhile, Emission rates ($E_{3km obs}$) at this 3 km grid of

observation site was shown. Fig. R2 and R3 showed the comparison of E_{45km} , E_{3km} and $E_{3km obs}$ for NO and C_2H_4 emission rates. Clearly, $E_{3km obs}$ at the observation sites were close to E_{45km} , which indicated that these observation sites rarely receive local emissions.

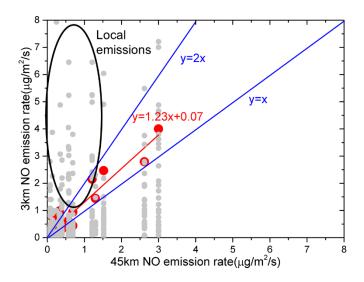


Fig. R2 Scatter plots of E_{45km}, E_{3km} and E_{3km obs} for NO emission rates (μg/m²/s). The gray and red solid cycles represented the relationships of E_{45km}-E_{3km} and E_{45km}- E_{3km obs}. The gray solid cycles above the line y=2x represented the grids receive the local emissions.

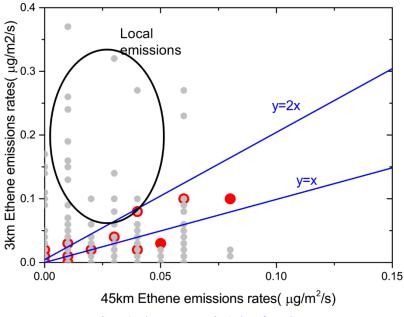


Fig. R3 The same as Fig.2, but for ethene

The following is a comment on the revised manuscript other than the above.

Comment 2: Page 11 Line 11: It describes the measuring instruments, but which measuring stations adopt these measuring instruments? **Reply:** In this study, NO_x was measured by Thermo Scientific 42C NO-NO₂-NO_x Analyzer with chemiluminescence technology at 40 sites in all three network (CERN, PRD-RAQMN). O₃ were measured by Thermo Scientific 49i with UV photometric technology from USA in CERN network and by Thermo Scientific 49C in PRD-RAQMN and EANET network.

Changes in the revised manuscript: Page 8 Line 14-17.

Comment 3: <About the response of my Comment 3>

It is said that the ensemble of models underestimates the measured values of dry deposition velocity from August to September in EA1, but according to Figure R3 it is actually overestimated.

Therefore, the claims in the response document or in the revised manuscript that underestimation of dry deposition velocity contributed to overestimation of summer ozone concentrations are false and need to be reviewed.

Reply: We agree. In the revised manuscript, we reworded related discussions and conclusions.

"In EA1, ensemble mean values overestimated observed dry deposition velocities of O₃ (vd) in August-September, but still fell into the range of observed standard deviation. This indicated that there must be other factors rather than dry deposition playing important roles in the overestimation of August-September O₃ in EA1. In October-November, simulated vd apparently underestimated observations by 30-50%. Among models, the lower dry deposition velocities in May-July from M1, M2, M4 and M6 than that of M11 partly explained higher May-July surface O3 from those simulations than that from M11. However, M13 and M14 still produced high O3 concentrations in May-September although their dry deposition velocities were similar to that of M11(Fig. 3)."

Changes in the revised manuscript: Page 17 Line 14-23.

The following are comments on the revised manuscript other than the above.

Comment 4: Section 3.5: The sentences are difficult to understand. I think that English native check is necessary first.

Reply: We invited an English native speaker to check our manuscript through AJE company.

Comment 5: Page 16 Line 4-5: I do not understand the meaning of this sentence. **Reply:** We revised to ". Quantifying the contributions of these processes can help explain model biases through sensitivity simulations"

Changes in the revised manuscript: Page 16 Line 24-25.

Comment 6: Page 16 Line 27 "summer": A statement from May to July, not summer,

is more logical. **Reply:** We agree and revised it.

Comment 7: Page 17 Line 3-4: According to Fig.9, contrary to the description in the revised manuscript, the model underestimated the observed values of Vd.
Reply: We agree and corrected it.
Changes in the revised manuscript: Page 16 Line 24-25.

Comment 8: Page 18 Line 7-8: The description in this sentence should indicate that it is a description of the observed value.

Reply: We agree. "The slope and intercept of regression line between observed O₃ and NOx were -0.77 ppbv/ppbv and 59.5 ppbv, respectively"

Changes in the revised manuscript: Page 18 Line 27.

Comment 9: Page 18 Line 16: What are "previous theoretical results"? **Reply:** According to comments by co-editor, we deleted these sentences.

Comment 10: Page 19 Line 15-25: At the beginning of chapter 5, you have mentioned that you evaluated dry deposition, PBL, and chemistry.

I do not understand why the discussion about the meteorological field comes out suddenly here.

Reply: In the revised manuscript, we remove this discussion.

Comment 11: <About the response of my Comment 6>

Page 5 Line 9-10: This sentence is about how to validate the model, so this sentence should be moved to chapter 3 which discusses it. **Reply:** We moved it to section 3.2.

Changes in the revised manuscript: Page 9 Line 25.

Comment 12: <About the response of my Comment 9>

Page 7 Line 2-6: Make a description that shows the correspondence between the model and the university.

Reply: We agree. "GEOS-Chem was run with a $2.5^{\circ} \times 2^{\circ}$ horizontal resolution and 47 vertical layers by University of Tennessee and Chemical AGCM for Study of Atmospheric Environment and Radiative Forcing (CHASER) was run with a $2.8^{\circ} \times 2.8^{\circ}$

horizontal resolution and 32 vertical layers by and Nagoya University."

Changes in the revised manuscript: Page 7 Line 6-8.

Comment 13: <About the response of my Comment 17>

It is understood from Appendix A that various statistics are derived taking into account only the variation due to the location (i) of the measuring station, but in fact it seems that temporal variations are also taken into consideration.

This is because, for example, in Page 11 Line 22-24, as the reason for the high correlation, the high reproducibility of the monthly variation is mentioned. **Reply:** We totally agree. We added the variation due to temporal variations in Appendix

Reply: We totally agree. We added the variation due to temporal variations in Appendix A.

Comment 14: <About the response of my Comment 27>

I understood that all models must be identified in Fig.2 and Fig.3. However, since it is still difficult to identify models by color, I hope that at least models that appear in the discussion in chapter 5 can be identified by such as drawn lines. **Reply**: We agree. In Fig. 3 and 4 in the revised manuscript, we only plotted models in appear in the discussion in chapter 5 by colors. The others are plotted in gray lines.

<Other technical corrections to the revised manuscript>

As there are many necessary technical corrections, I would like you to carefully review the entire manuscript.

Some of them are shown below.

Comment 15: Page 2 Line 2: "Evaluated and intercompared to O3 observations" perhaps should be "intercompared and evaluated to O3 observations". **Reply:** We revised it.

Comment 16: page 2 Line 7-8: "western pacific rim" is repeated twice. **Reply:** We deleted the second one.

Comment 17: Page 8 Line 25: "Table1" should be "Table2". Reply: We revised it.

Comment 18: Page 11 Line 27: "Table1" should be "Table2".

Reply: We revised it.

Comment 19: page 14 line 27: The word "other" should be added before "combined influence". **Reply:** We revised it.

Reply: we levised it.

Comment 20: Page 17 Line 24 "Sillman and He et al.": "Et al" should be removed. **Reply:** We revised it.

Comment 21: Page 20 Line 10-11: There is an incomplete sentence.

Some of figures: "EA3" should be replaced by "EA2", and "EA4" should be replaced by "EA3".

Reply: We revised it. "For the North China Plain and western Pacific Rim, the model ensemble severely overestimated surface O3 levels for May-September by 10-30 ppbv."

Changes in the revised manuscript: Page 23-25

We thank co-editor for his constructive comments. Response to the Specific comments.

General comment: Although the manuscript is improved, still many fundamental and technical points are present needing revision. Please examine comments from the reviewer #1 carefully, and the following comments by the Co-Editor. Recheck English throughout the manuscript with a native speaker.

Reply: Thanks a lot for insightful comments. In the revised manuscript, we invited a native speaker by AJE company to improve our language in this manuscript.

(pages and lines are for the change-track version)

Comment 1: Page 2, line 18: In terms of **Reply:** We revised it.

Comment 2: Page 2, line 20: hPa **Reply:** We revised it.

Comment 3: Page 3, line 21: do not **Reply:** We revised it.

Comment 4: Page 5, lines 19-20: Better to place the following inserted sentence elsewhere (page 8-9?): If two or more observation sites were in the same grid of model, their mean values will be used to evaluate model performance. **Reply:** We placed it into Page 9 Line 25.

Comment 5: Page 6, line 18. possibly **Reply:** We revised it.

Comment 6: Page 8, lines 18-20: More explanation and clarification about the local sources and sinks are necessary, as requested by the reviewer. Not only for NOx but also VOC, representativeness is necessary.

Reply: We agree. In the revised manuscript, we plotted NO and VOCs emission rates.

Before the intercomparison between emission rates in 45 km and 3 km emission inventory, we provide the equation how to calculate the emission rates ($\mu g/m^2/s$) in 45 km resolution (E_{45km}) from those in 3 km resolution (E_{3km}).

 $E_{45km} = \sum_{i=1}^{m=15} \sum_{j=1}^{n=15} (E_{3km\,ij} / (m \times n)) \quad (1)$

As shown in Eq (1), some $E_{3km ij}$ were higher than E_{45km} . Others could be less than E_{45km} . In general, there could be local emissions in ith and jth grid when E_{3km} in this grid is much higher than E_{45km} . When E_{3km} is much less than E_{45km} , the ith and jth grid represented a regional background condition. If E_{3km} is close to E_{45km} , the emission

rates in the ith and jth grid represented the averaged conditions of 45 km \times 45 km areas. So, the comparison of E_{3km} with E_{45km} is a good way to examine if this 3km grid receives local emissions.

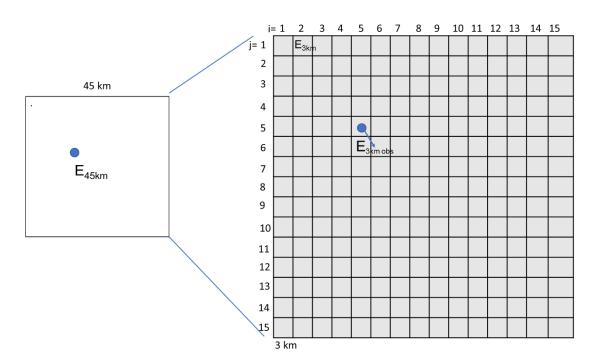


Fig. R1 The grids of observations in 45 (left) and 3 km (right) emissions inventory. The solid cycle represents the location of observation site.

In this study, we calculated the NO emission rates of observation sites in 45 km grid (E_{45km}). Emission rates (E_{3km}) at each 3 km grid (i=1,15; j=1,15) within this 45km grid was compared with E_{45km} . Meanwhile, Emission rates ($E_{3km obs}$) at this 3 km grid of observation site was shown. Fig. R2 and R3 showed the comparison of E_{45km} , E_{3km} and $E_{3km obs}$ for NO and C_2H_4 emission rates. Clearly, $E_{3km obs}$ at the observation sites were close to E_{45km} , which indicated that these observation sites rarely receive local emissions.

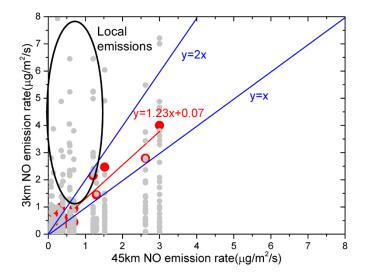
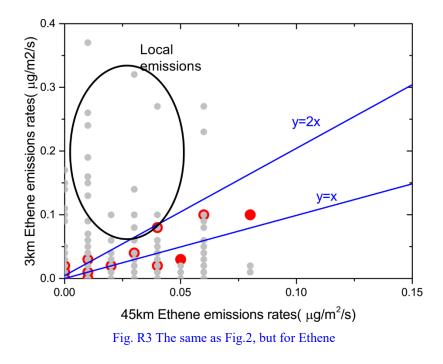


Fig. R2 Scatter plots of E_{45km}, E_{3km} and E_{3km obs} for NO emission rates (μg/m²/s). The gray and red solid cycles represented the relationships of E_{45km}-E_{3km} and E_{45km}- E_{3km obs}. The gray solid cycles above the line y=2x represented the grids receive the local emissions.



Comment 8: Page 8, line 28: exhibited **Reply:** We revised it.

Comment 9 Page 9, line 1: Molybdenum converters **Reply:** We revised it.

Comment 10: Page 9, lines 2-5. The results must be dependent on the conditions (with different NOz/NOx). Where did the comparison take place? Uncertainty should be estimated to cover all possible conditions (urban to rural) used for the study.

Reply: We agree. In this study, the comparison was conducted in an urban site in Beijing from 1 to 29 August 2012. As said by the editor, the difference between NO₂ and NO₂* is dependent on the conditions (Fig. R1). In general, their differences appeared within 10-15% when NO₂*>15 ppbv and NOz<10 ppbv in Beijing. In low NO₂ (<15 ppbv) and high NOz (>10 ppbv) conditions, the NO₂* usually was higher than NO₂ (10-30%). Unfortunately, there are very few CAPS NO₂ measurements in China. Jung reported their NO₂ and NO₂* usually appeared within 20% when NO₂* >20 ppbv in Spring and Summer. NO₂* overestimated NO₂ by 20-40% when NO₂* <20 ppbv. In Fall and Winter, NO₂* usually overestimated NO₂ by 10-20 ppbv in all conditions.

In the revised manuscript, we added a short discussion on NO₂* uncertainty.

"This bias was dependent on the chemical conditions. A one-month continuous measurement in August by a chemiluminescence analyzer and Aerodyne Cavity Attenuated Phase Shift Spectroscopy (CAPS) at an urban site in Beijing showed that this bias from a chemiluminescence analyzer was small when NO2 concentrations were more than 10-15 ppbv, and ranged from 10% to 30% under low NO2 (<10 ppbv) (Ge et al., 2013). Measurements at a rural site in South Korea also revealed a similar pattern (Jung et al., 2017). These comparisons suggested that observations by molybdenum converters may overestimated NO2 by 10-20% in EA1 and EA2, and 30% in EA3. This brings uncertainties for the model evaluation on NO₂ in this study. "

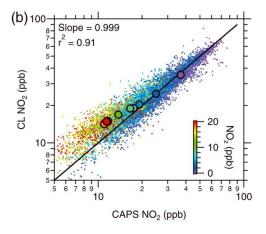


Fig. R1 Comparison of hourly NO₂ measured by the CL NO_x analyzer and the CAPS NO₂ monitor in Beijing

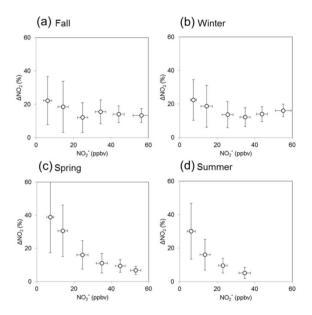


Fig. R2 The NO₂/NO₂* ratios as a function of NO₂ * concentrations during (a) fall, (b) winter, (c) spring, and (d) summer in a rural site in South Korea (Jung et al., 2017). NO₂ and NO₂ * is the measurements by the photolytic converterver and molybdenum converter .

Changes in the revised manuscript: Page 8 Line 20-27.

Comment 11: Page 9, line 16: parts per billion by volume **Reply:** We revised it.

Comment 12: Page 9, line 19: and M7 **Reply:** We revised it.

Comment 13: Page 9, line 20: closer to **Reply:** We revised it.

Comment 14: Page 9, line 20: M11 simulated O_3 with RMSEs of **Reply:** We revised it.

Comment 15: Page 12, line 26: existed **Reply:** We revised it.

Comment 16: Page 13, line 21. approximately 20 and 10 ppbv **Reply:** We revised it.

Comment 17: Page 14, line 7: showed a better performance

Reply: We revised it.

Comment 18: Page 15, line 5: What is meant by short-lived species? **Reply:** This underestimation partly was attributed to the coarse model horizontal resolution (45km) used in the MICS-Asia III, which hardly reproduced concentrations of short-lived species (e.g. NO).

Comment 19: Page 15, line 25: Figure 5 shows **Reply:** We revised it.

Comment 20: Page 15, lines 27-28 and elsewhere: hPa **Reply:** We revised it.

Comment 21: Page 16, line 7: The mean standard deviations (SD) of models (1σ) Use SD instead of σ later. **Reply:** We revised it.

Comment 22: Page 23, line 16. MICS **Reply:** We revised it.

Comment 23: Page 27, lines 23. Ensemble mean values rather OVERESTIMATES vd (Fig. 9). Discussion needs to be revised (conclusion also). Representativeness of deposition velocity observations over the domains needs to be discussed. I doubt that grassland represents the whole domain region.

Reply: We agree. We reworded discussions (and conclusions) on dry depositions in the revised manuscript.

"In EA1, ensemble mean values overestimated observed dry deposition velocities of O3 (vd) in August-September, but still fell into the range of observed standard deviation. This indicated that there must be other factors rather than dry deposition playing important roles in the overestimation of August-September O₃ in EA1. In October-November, simulated vd apparently underestimated observations by 30-50%. Among models, the lower dry deposition velocities in May-July from M1, M2, M4 and M6 than that of M11 partly explained higher May-July surface O₃ from those simulations than that from M11. However, M13 and M14 still produced high O₃ concentrations in May-September although their dry deposition velocities were similar to that of M11."

In this study, we selected observations at two sites in EA1. One site was located in a valley of Mangshan Forest Park, the other (CREAS) was in a suburb about 10km north of the Beijing city. The CREAS site had an area of $200m \times 200m$, and was thickly covered with short grass about 10cm high (Dense grassland). Fig.R3 presents the land cover classification map in EA1 from MODIS satellite data (Zhang et al., 2008). It can be found that dense grassland is one of the dominated landcover classes and covers ~20% land area. Another dominated landuse in EA1 is crop class, which covers more

than 50% land area. Unfortunately, there are few observations on O_3 dry deposition on the crop class in China. This may bring uncertainties for model evaluation. Hardacre et al. (2015) reported O_3 dry deposition measurements on crops in Europe and simulated O_3 dry deposition in 15 global models. Both observations and simulations showed that O_3 dry deposition velocities on agriculture crop class were quite similar as grasslands (Fig. R4). This indicated that the uncertainties on representativeness of measurement sites in this study did not affect our conclusions.

We added related discussions in the revised manuscript.

Changes in the revised manuscript: Page17 Line 25-Page 18 Line 2.

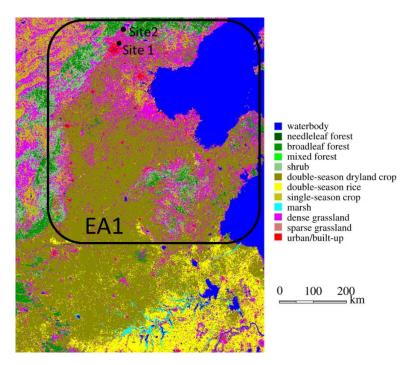


Fig. R3 The land cover classification map from MODIS satellite data

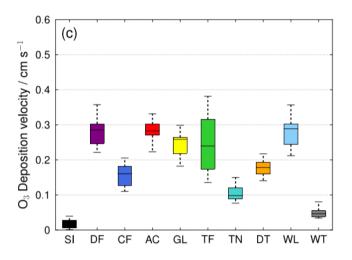


Fig. R4 Total annual O₃ dry deposition and annual average O₃ deposition velocity partitioned to land cover classes (GL: grassland; AC: agriculture crop class)

Comment 24: . Page 28, lines 3-4. October-November, simulated vd apparently overestimated observations by 30-50%. This is also opposite.

Reply: We agree. In the revised manuscript, we corrected it. "In October-November, simulated vd apparently underestimated observations by 30-50%."

Changes in the revised manuscript: Page 17 Line 18.

Comment 25: Page 29, line 6. Is this only because O3 is titrated with NO? The analysis never tells the regimes. Consider cases of O3 buildup and transport by consuming NOx. Even in the NOx limited regime, the NOx-O3 relationship constructed from geographical distribution will show a negative correlation. At least potential ozone ($O_3 + NO_2$) should be used for the analysis, and remove all statements on the regimes (particularly those in lines 22-23, page 29).

Reply: We agree. In the revised manuscript, we remove all statements on the regimes.

Fig. R5 shows the relationship between NO_x and O_x(O₃+NO₂) in this study. For EA1, observed O_x increases with the increase of NO_x levels, with coefficient of determination (R^2) of 0.61. Most of the models (except for M8, M11 and M13) failed to reproduced observed positive correlations between O_x and NO_x, and their R^2 only ranged from 0.01-0.08. The slope, intercept and R^2 of M8 and M11 are relative agreement with observations. For EA2, all models reproduced observed key patterns in which O_x positively correlated with NO_x.

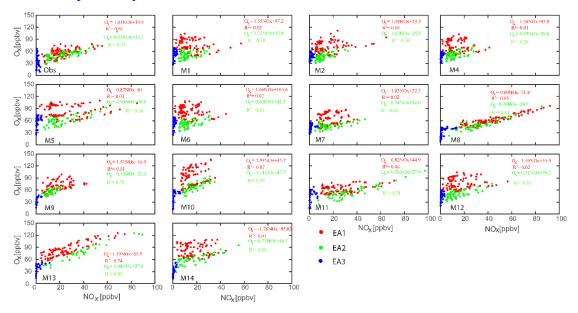


Fig. R5 Scatter plots for monthly daytime (08:00-20:00) surface NO_x and O_x for each station in EA1 (red), EA2 (green)and EA3 (blue) in May-October, for observations (obs) and models. Also shown are the linear regression equations and coefficient of determination (R^2) for NO_x and O_x (O_3+NO_2) in EA1 (red) and EA2 (green).

Changes in the revised manuscript: Page 19 Line 7-15.

Comment 26: Page 30, line 6. Briefly mention how PBLH was measured.

Reply: In this study, the observed PBLH was derived from the radiosonde network of the L-band sounding system of the China Meteorological Administration (CMA). The system provides fine-resolution profiles of temperature, pressure relative humidity, wind speed and direction. the bulk Richardson number (Ri) method (Vogelezang and Holtslag,1996) was taken to simultaneously estimate the PBLH from CMA soundings. R_i is defined as the ratio of turbulence associated with buoyancy to that induced by mechanical shear, which is expressed as

$$Ri(z) = \frac{(g/\theta_{vs})(\theta_{vz} - \theta_{vs})(z - z_s)}{(u_z - u_s)^2 + (v_z - u_s)^2 + (bu_*^2)},$$

where z denotes height above ground, s the surface, g the acceleration due to gravity, θv virtual potential temperature, u and v the component of wind speed and u* the surface friction velocity. u* can be ignored here due to the much smaller magnitude compared with bulk wind shear term in the denominator (Vogelezang and Holtslag, 1996). The critical value of 0.25 (Ri) is referred to as PBLH in this study, similar to the criteria used by Seidel et al. (2012).

Changes in the revised manuscript: Page 19 Line 27-Page 20 Line 5.

Comment 27: Page 30, line 8. caused by the sampling bias between **Reply:** We revised it.

Comment 28: Page 31, line 9. A model ensemble was produced **Reply:** We revised it.

Comment 29: Page 31, line 14. A period character must be a comma? **Reply:** We revised it.

"In North China Plain and western Pacific rim, the model ensemble severely overestimated surface O₃ in May-September by 10-30 ppbv."

Comment 30: At many times figures contain EA1, 3, and 4, instead of EA1, 2, and 3. **Reply:** We corrected it.

- Jinsang Jung, JaeYong Lee, ByungMoon Kim, SangHyub Oh, Seasonal variations in the NO2 artifact from chemiluminescence measurements with a molybdenum converter at a suburban site in Korea (downwind of the Asian continental outflow) during 2015–2016, Atmospheric Environment, Volume 165,2017,Pages 290-300.
- Seidel, D. J., Zhang, Y., Beljaars, A., Golaz, J.-C., Jacobson, A. R., and Medeiros, B.: Climatology of the planetary boundary layer over the continental United States and Europe, J.Geophys.Res.Atmos., 117, D17106, doi:10.1029/2012JD018143, 2012
- Vogelezang, D. H. P. and Holtslag, A. A. M.: Evaluation and model impacts of alternative boundary-layer height for mulations, Bound.-Lay. Meteorol., 81, 245–

269, doi:10.1007/BF02430331, 1996.

- Xia Zhang, Rui Sun, Bing Zhang, Qingxi Tong, Land cover classification of the North China Plain using MODIS_EVI time series, ISPRS Journal of Photogrammetry and Remote Sensing, Volume 63, Issue 4,2008,Pages 476-484.
- Hardacre, C., Wild, O., and Emberson, L.: An evaluation of ozone dry deposition in global scale chemistry climate models, Atmos. Chem. Phys., 15, 6419-6436, https://doi.org/10.5194/acp-15-6419-2015, 2015.

Model evaluation and inter-comparison of surface-level ozone and relevant species in East Asia in the context of MICS-Asia phase III Part I: overview

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Abstract: Long-term ozone (O_3) and nitrogen oxide (NO_x) from fourteen state-of-the-art chemical transport models (CTMs) are intercompared and evaluated evaluated and intercompared to with O₃ observations in East Asia, within the framework of the Model Inter-Comparison Study for Asia phase III (MICS-ASIA III), designed to evaluate the capabilities and uncertainties of current CTMs simulations in for Asia and to provide multi-model estimates of pollutant distributions. These models were run by fourteen independent groups working in China, Japan, South Korea, the United States and other countries/regions. Compared with to MICS-Asia II, the evaluation against of observations was extended to be-one-full year in-across China and the western Pacific Rim from four months-and the western Pacific Rim. In general, the model performance levels for O_3 varied widely, depending on by region and seasons. Most models captured the key patterns of monthly and diurnal variation of surface O_3 and its precursors in North China Plain and western Pacific Rim, but failed to do so for the in-Pearl River Delta. A significant overestimation of surface O₃ was evident in-from May-September/October and from January-May over the North China Plain, western Pacific Rim and Pearl River Delta. Comparisons with-drawn from observations revealed show that underestimation on dry deposition velocities and largeconsiderable diversity of in photochemical production partly contributed to this overestimation and large to high levels of intermodel variability ion O_3 form North China. In terms of O_3 soundings, the ensemble average of the models reproduced the vertical structure in-for the western Pacific, but overestimated O₃ levels to below 800 hpPa in the summer. In the industrialized Pearl River Delta, the ensemble average presented an overestimation in for the lower troposphere and an underestimation in for the middle troposphere. The ensemble average of 13 models for O₃ did not always exhibit a-superior performance compared to certain

ensemble average of 13 models for O₃ did not always exhibit a-superior performance compared to certain individual models₅ in contrast to its superior value fority in Europe. This finding suggestsed that the spread of ensemble-model values had-does not represented all uncertainties of O₃ or that most models in MICS-Asia III models missed key processes. This study improves d-the performance of modeling O₃ in March at Japanese sites than the previous phase of MICS-Asia ((MICS-Asia II)). However, it
 overpredicteoverpredictsd surface O₃ concentrations in for western Japan in July, which has was not been found in by MICS-Asia II. Major challenges still remain in regard to identifying the sources of bias in

surface O3 over East Asia in CTMs.

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1. Introduction:

Tropospheric ozone (O₃) is a significant secondary air pollutant produced through thousands of photochemical reactions and <u>that is</u> detrimental to human health, ecosystems, and climate change as a strong oxidant (WHO, 2005; The Royal Society, 2008). With <u>rapid the fast</u>-industrialization and urbanization in the last two decades, O₃ concentration is rising at a higher rate in East Asia than <u>in</u> other regions_a and <u>on</u>_30% of the days in megacities (e.g. Beijing, Shanghai Guangzhou in China) <u>values</u> exceeds air quality standard of World Health Organization (100 μ g/m³) for 8-hour average surface O₃ concentration (Wang et al.,2017). <u>TheH-high</u> O₃ concentrations <u>have</u> received more attention from the public and from policy-makers in East Asia. The Ministry of Environment Japan has imposed stringent measures to reduce traffic emissions since <u>the</u>_1990s, and non-methane volatile organic compounds (NMVOCs) and NO_x mixing ratios have decreased by 40-50-% and 51-54-%, respectively (Akimoto et al.,2015). In 2012, China released a new ambient air quality standard <u>in-under</u> which <u>a limit on</u> the 8-hour O₃ maximum was set <u>limits</u>-for the first time. However, these measures <u>don't do not</u> prevent the persistent increase of the ground-level O₃ in East Asia. The average mixing ratio of O₃ <u>has</u> increased 20-30% in Japan over the last 20 years (Akimoto et al.,2015). In Chinese megacities, 8-hr O₃ concentrations have increased 10-30-% since 2013 (Wang et al.,2017).

The primary-main method used for the detailed evaluation of the effects of air quality policies at the scale of East Asia is that of numerical air quality modeling. Several global and regional scale CTMs (e.g. GEOS-Chem, CHASER, CMAQ, CAMx, WRF-Chem and NAQPMS) have been developed over the past few decades have been developed and have been widely used to simulate the O₃ formation process and to evaluate strategies for its control strategies-(Streets et al., 2008; Li et al., 2007; 2008; Yamaji et al., 2006; Zhang et al., 2008; Liu et al., 2010; Wang et al., 2013; He et al., 2017; Nagashima et al., 2017). These-Such simulations have identified the key precursors of O₃ formation in East Asia (Zhang et al., 2008; Liu et al., 2011; He et al., 2017), have assessed the contributions of international and regional transport (Streets et al., 2008; Li et al., 2008), and have predicted the-O₃ mixing ratios underim different future emission scenarios (Wang et al., 2013). However, discrepancies remain between models and observations, indicating that model simulations of O₃ in East Asia still need to be improved (Han et al., 2008). Modeling uncertainties related to the emissions, chemistry, wet and

dry deposition, and transport can hardly be handled-addressed using one a single model. Model intercomparison has thus been recognized as an effective way to address problems and has been successfully applied in Europe and North America in the phase 2 of the Air Quality Model Evaluation International Initiative (AQME II; Rao et al., 2011). Limited model inter-comparison related to air quality in East Asia has been conducted. Phases I and II of the Model Inter-Comparison Study for Asia (MICS-Asia) were initiated in 1998 and 2003, and to explore the potential sources of model uncertainties regarding sulfur, O₃, nitrogen compounds and aerosols (Carmichael et al., 2002, 2008). They found study shows that the predicted temporal variations of found for surface O₃ in eight regional CTMs generally tended to be lower than that those observed in 2001 with poor correlations in the western Pacific in March and December (Han et al., 2008). Model performance levels for O₃ were found to varyied largely greatly in southern China. The iInconsistenciesy iof horizontal grids, emissions and meteorological inputs used among models increased the difficulty of explaininghave rendering explaining intermodel variability in theusing MICS-Asia II results more difficult. More importantly, model evaluations form industrialized China haves not been conducted because of due to a lack of few observations, which has been detrimental to efforts made to improve O₃model performance levels on O₃.

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Recently, regional CTMs have been greatly improved by coupling more mechanisms (e.g. heterogeneous chemistry and on-line calculation of photolysis rates) and accurate chemical reaction rates. For example, the-gas-phase chemistry mechanisms in of Models 3-Community Multiscale Air Quality (CMAQ) have been developed into CBM05 and SAPRC07 from CB04 and SAPRC99. It is critical to evaluate the updated models' abilities for to simulateing current air quality levels over East Asia. In 2010, MICS-Asia was expanded to Phase III wherein, in which-13 regional CTMs and 1 global CTM are were run over one-full year by 14 independent groups from East Asia and North America, under-using a common reference model input data set (namely, the emission inventory, meteorological fields and horizontal grids). In addition to observations made in Japan by the Acid Deposition Monitoring Network in East Asia (EANET) that were used in-for MICS-Asia II, new observational data from China were made available for MICS-Asia III and, which were obtained from the Chinese Ecosystem Research Network (CERN) and the Pearl River Delta Regional Air Quality Monitoring Network (PRD RAQMN). An intercomparison of CTMs in China, Japan and the western Pacific over-for one full year have never

<u>before</u> been performed, <u>which provided_creating</u> a <u>broader</u> <u>wider</u>-database to use <u>forin the</u> comparisons. The completeness of MICS-Asia III is therefore unique.

In this paper, we mainly evaluate the <u>abilities capacities</u> of participating models in MICS-Asia III for to simulate the concentrations of O_3 and its related species in the framework within of the MICS-

5 Asia III framework. Several-The follwing questions are addressed: (1) What is the performance level ofHow well do various air quality models for-perform in simulating O₃ levels in East Asia? (2) How consistent or discrepant are the models? (3) How do muti-model ensembles improve O₃the simulation accuracy for O₃? This paper is expected to provide valuable insights into the abilities-capacities and limitations of CTMs when applied to in East Asia.

10 2. Models and data

2.1 Experimental set up

In this study, all participating models were run for the year 2010 and provide gridded monthly mean diurnal O_3 and its precursors mixing ratios in the lowest model layer. For O_3 , monthly three-dimensional data were also submitted. If two or more observation sites were in the same grid of model, their mean values will be used to evaluate model performance.

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2.2 Participating models and input data

Table 1 summarizes the specifications of participating CTMs. These models include two versions of CMAQ (v4.7.1 and 5.0.2; Byun and Schere, 2006), the Weather Research and Forecasting model 20 coupled with Chemistry (WRF-Chem; http://www.acd.ucar.edu/wrf-chem), the Nested Air Quality Prediction Modeling System (NAQPMS; Li et al., 2007), the Japan Meteorological Agency (JMA)'s nonhydrostatic meteorology-chemistry model (NHM-Chem; Kajino et al., 2012), the NASA-Unified Weather Research and Forecasting (NU-WRF; Tao et al.,2013) and **GEOS-Chem** (http://acmg.seas.harvard.edu/geos/). They have been documented in the scientific literature and have 25 been widely applied in modeling studies over of East Asia. Table 1 did does not list model names to maintain each model's anonymity for each participating model. Similar behavior was also found observed fromin MICS-Asia II and other model intercomparison projects (e.g., AQME II).

MICS-Asia III participants were provided with a reference meteorological field for the year 2010, generated with the Weather Research and Forecasting Model (WRF) version 3.4.1 model. The domain of meteorological fields is shown in Fig. 1. WRF v3.4.1 are driven by the final analyses dataset (ds083.2) from the National Centers for Environmental Prediction (NCEP), with $1^{\circ} \times 1^{\circ}$ resolution and a temporal

resolution of 6 h. A four-dimensional data assimilation nudging toward the NCEP dataset was performed to increase the accuracy of the_WRF. The horizontal model domain_of, which is __182 ×172 grids on a Lambert conformal map projection with 45-km horizontal resolution, is shown in Fig. 1. Vertically, the WRF grid structure consists of 40 layers from the surface to the top of model top-(10 hPa.). The sStandard meteorological fields were applied by the majority of groups. Several other models were employed to performed simulations using their own meteorological models (e.g., RAMS-CMAQ and GEOS-Chem). The WRF-Chem utilized the same model (WRF) as the standard meteorological simulation; but they considered the feedback of pollutants to the-meteorological fields. Consequently, their meteorological fields are possible-may be slightly different from the standard. GEOS-Chem is driven by the GEOS-5 assimilated meteorological fields taken from the Goddard Earth Observing System of the NASA Global
Modeling Assimilation Office. The couples of meteorological data and CTMs vary for each group, likely

resulting in a diversified set of model outputs.

MICS-Asia III provided-provides a set of monthly anthropogenic emission inventory-inventories for the year 2010, which is-called as-MIX (Li et al., 2016). MIX is a mosaic of up-to-date regional and national emission inventories that includes Regional Emission inventory in ASia (REAS) version 2.1 for
the whole of-Asian region (Kurokawa et al., 2013), the Multi-resolution Emission Inventory for China (MEIC) developed by Tsinghua University, a high-resolution NH₃ emission inventory by Peking University (Huang et al., 2012), an Indian emission inventory developed by Argonne National Laboratory (ANL-India, Lu et al., 2 011; Lu and Streets, 2012), and the official Korean emission inventory from the Clean Air Policy Support System (CAPSS; Lee et al., 2011). BThe biogenic emissions are taken from the Model of Emissions of Gases and Aerosols from Nature (MEGAN). Hourly biogenic emissions were obtained for the entire year of 2010 using version 2.04 (Guenther et al., 2006). Biomass burning emissions were processed by re-gridding the–Global Fire Emissions Database version 3 (GFEDv3) (0.5 by 0.5 degree). Volcano SO₂ emissions were provided, with a daily temporal resolution

by the Asia Center for Air Pollution Research (ACAP). The emission group in-MICS-ASIA III emission group_directly prepared a gridded inventory according to the configuration of each CTM. NMVOC emissions are spectated into model-ready inputs for three chemical mechanisms: CBMZ, CB05 and SAPRC-99. Weekly and diurnal profiles were also provided. The standard emission inventory was applied by all models. The majority of models employed official suggested vertical and time profiles of pollutants from each sector by the emission group. M13 and M14 maked the projections by themselves.

MICS-Asia III also provided two sets of chemical concentrations <u>forat</u> the top and lateral boundaries of the model domain, which were derived from the-3-hourly global model outputs for the year 2010. The

More information can be found in the paper of Li et al. (2017) and Gao et al. (2017).

global models were run by University of Tennessee (http://acmg.seas.harvard.edu/geos/) and Nagoya
 University (Sudo et al., 2002). GEOS-Chem was run with a 2.5°×2° horizontal resolution and 47 vertical
 layers by University of Tennessee, and Chemical AGCM for Study of Atmospheric Environment and
 Radiative Forcing (CHASER) was run with a 2.8°× 2.8° horizontal resolution and with 32 vertical layers
 by Nagoya -University. Some models made-applied boundary conditions depending on their own
 previous past_experiences.

2.3 Observational data for O₃

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In this study, East Asia has been<u>was</u> divided into three sub-regions as shown in Fig. 1. The selection of the sub-regions <u>wasis</u> based on emissions, climate and observation data coverage. The North China Plain (EA1) and Pearl River Delta (EA2) represent the highly industrialized regions <u>of</u> the mid-latitudes. EA1 <u>is characterized byhave</u> a temperate and tropical continental monsoon climate with marked seasonality, respectively. EA2 is located in the south of<u>ern</u> China, and is less affected by the continental air masses. EA3 <u>consists-covers</u> of the northwest Pacific and the Sea of Japan, and represents the downwind regions of the Asian continent with a marine climate.

Hourly O₃ and NO_x observations in-for the year 2010 in East Asia were obtained from the CERN,
PRD-RAQMN), and EANET. The CERN was built by the Institute of Atmospheric Physics, Chinese Academy of Sciences and consists of includes 19 surface stations within covering an area of 500 × 500 km² in-across North China Plain (EA1 sub-region; Ji et al., 2012). The se-stations were set up according to the United States Environmental Protection Agency method designations. Half of them were

remote, rural, or-suburban and clear urban sites. Nine9 sites were located in the within meteorological stations or on campuses of universities in urban regions, with little influence from local sources and sinks. The A comparison of NO and ethene emission rates at these sites in 45 km and 3 km resolution emission inventories showeds that observations generally represented the ~45 km averages of pollutants. The PRD RAQMN was jointly established by the governments of the Guangdong Province and the Hong Kong Special Administrative Region and consists of 16 automatic air quality monitoring stations located across the EA2 sub-region (Zhong et al., 2013). Thirteen of these stations are operated by the Environmental Monitoring Centers in Guangdong Province and the other three are located in Hong Kong and are managed by the Hong Kong Environmental Pollution Department. The PRD RAQMN was designed to probe the regional air quality, to assess the effectiveness of emission reduction measures and to enhance the roles of monitoring networks in characterizing regional air quality and in supporting air quality management. SoThus, the sites are rarely influenced by local sources and sinks. The EANET was launched in 1998 to address acid deposition problems in East Asia, following the model of the Cooperative Program for Monitoring and Evaluation of the Long-Range Transmission of Air Pollutants in Europe. In this study, eight remote stations in the northwestern -Pacific and Japan (EA3 sub-region) were selected for use in to evaluate ing the model performance levels in for the downwind regions of the Asian continent (Ban et al., 2016). More information on the EANET can be found at http://www.eanet.asia/. Note that only stations with at least 75% data validity were chosen. Table S1 in the supplementary sections lists provides detailed site description. Our comparisons of NO_x and VOCsemission rates conducted on grid for these stations at 45 km and 3 km resolution emission inventories suggest that our selected stations have rarely received local emissions.

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 O_3 were measured by Thermo Scientific 49i with UV photometric technology in CERN network and by Thermo Scientific 49C in PRD-RAQMN and EANET network. NO_x was measured by Thermo Scientific 42C NO-NO2-NOx Analyzer with chemiluminescence technology at 40 sites in all three

25 <u>networks (CERN, PRD-RAQMN and EANET).</u> The O₃- and NO_x-instruments were an ultraviolet photometric analyzer (model49i, Thermo Fisher Scientific (Thermo), USA) and a chemiluminescence analyzer (model42iTL, Thermo, USA), respectively. NO_x measurement<u>s</u> existed <u>exhibited</u> sometimes biases (especially for stations <u>located</u> far from sources) when using <u>molybdenum converter molybden</u> eonvertors devices since all nitrogen oxydes are-were measured. <u>This bias was found to be dependent on</u> <u>chemical conditions.</u> A one-month continuous measurement <u>collected</u> in August <u>by with a</u> chemiluminescence analyzer and Aerodyne Cavity Attenuated Phase Shift Spectroscopy (CAPS) <u>from</u> an urban site in Beijing showed shows that this bias from athe chemiluminescence analyzer was is minor small-when NO₂ concentrations were more than<u>exceed</u> 10-15 ppbv, and ranginged from 10% to 30%

under low NO₂ <u>conditions (<10 ppbv)</u> (Ge et al., 2013). <u>Measurements collected from a rural site in South</u> <u>Korea reveal a similar pattern across all seasons (Jung et al., 2017)</u>. These comparisons suggest that <u>observations made using molybdenum converters may overestimate NO₂ by 10-20% for EA1 and EA2</u> and 30% for EA3, introducing uncertainties into the NO₂ model evaluation in this study.

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3. Model validation and general statistics

3.1 Annual concentrations of surface O₃, nitric oxide (NO) and nitrogen dioxide (NO₂)

Fig. 2 provides a concise comparison of model performances on-for annual O₃, NO and NO₂ in-for three sub regions in of East Asia. A box-and-whisker representation was used to show the frequency distribution of monthly concentrations measured fromat stations in each sub-region. The O₃ normalized mean bias (NMB) and root mean square error (RMSE) of the ensemble mean were found to be significantly less than the ensemble median in most csituations-ases (Table 12). Therefore, we only presented the results of multi-model mean ensemble results (Ense). In general, the majority of the models significantly overestimated annual surface O₃ relative to compared with the observations of in EA1, EA2 and EA3. Ense overestimated surface O₃ by 10-30 parts per billion by volume (ppbv) forin these subregions. Ense NO₂ levels was-closely_-reflect to the observations to within ±20% acrossin all subregions. In EA1 and EA2, Ense NO levels were found to be was-5-10 ppbv lower than those observed observation, while it showed exhibiting a-reasonable_-performance inlevels for EA3.

Among all<u>Of the</u> models, M11 in-for sub-regions EA1 and EA2,-and M7 in for EA2 and EA3 more 25 closely reflect were closer-O₃ observations. M11 simulated O₃ achieved with RMSEs of 9.5 ppbv and 13.3 ppbv in-for EA1 and EA2, respectively (Table 2). The performance levels of models' performance in for simulating O₃ were was found to be closely related to their performances for NO₂ and NO. In

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highly polluted regions (EA1 and EA2), a persistent underestimation of NO was evident across most models. A<u>s</u> an interesting phenomenon, we found the was that models' performance regarding O₃ <u>measurements to variedy</u> greatly for in EA3, although they but M8 showed exhibited a a consistent performance with respect to NO and NO₂. This finding suggests that O₃ was significantly affected by other factors in addition to local chemistry in EA3. M8 underestimated O₃ and overestimated NO in all sub-regions by 40-50%. The strongest highest O₃ titration level observed in M8 may have result generate in lower O₃ levels than those indicated by other models and observations.

3.2 Monthly variation of surface O₃, NO and NO₂

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_____Fig. 3 presents the monthly mean concentrations of O₃, NO and NO₂ in <u>for the three sub-</u> 10 regions <u>over across East Asia</u>. <u>When If two or more observation sites are located were in the same model</u> <u>grid-of model</u>, their mean values <u>willare be-used to evaluate model performance</u>.

All models captured the observed seasonal cycles of O₃, NO and NO₂ in-for_EA1. In-From May-September, Ense O_3 was 10-30 ppbv higher than observed values observations, (30-70% of observed values), while Ense NO and NO₂ levels appeared to be consistent with observations, attaining with mean 15 biases of < 3 ppbv. This finding suggests that anthe intercomparison of O_3 production efficiency levels per NO_x with observations is needed. In-For EA2, Ense O₃ values agreed well with observed high autumn O_3 levels, but are overestimated from January to September by 5-15 ppbv (15-60% of observations). This overestimation reached the highest point from in March-April (15ppbv) and led to a spring peak in simulated O₃ values which was not found in the observations. This overestimation iwas partly related to 20 the underestimation of NO in the same months, which decreased the titration effect. For NO₂, Ense value agreed well with observed values forin June-December, and slightly underestimated observations forin January-May. ForIn EA3, the ensemble NO₂ was generally close to the observed values to within ± 0.5 ppbv. A <u>S</u>ignificant overestimations of O₃ and underestimations of NO were observed from existed in June-October. Similar results have been found fromin MICS-Asia II and through another model inter-25 comparison project under-of the Task Force on Hemispheric Transport of Air Pollution (TF HTAP), which suggestinged that such results may stem from the differences in the representations of dispersion by southwesterly clean marine air masses dispersion observed across in different metrological fields used in CTMs (Han et al., 2008; Fiore et al., 2009).

For individual models, M11 achieved the best-highest degree of model reproductivity foref monthly mean O₃ levels in EA1 among of the examined models. Most of the Oother most-models overestimated O₃ by 100-200% form May-October. The largest levels of model bias and intermodel variability for NO and NO₂ appeared in the winter, which and likely came from the NO_x vertical diffusion and heterogeneous chemistry (Akimoto et al., 2019). In EA2, M7 seems to have achieved the best-highest levels of O₃ reproducibility for O₃. Most of the models (except for M11 and M12) exhibited show high O₃ concentrations form March-May and September-November. Observed O₃ values showed that the highest concentrations appeared from October-November. M11 captured the observed January-May O₃ value because of due to relatively high NO concentrations. However, NO was overestimated by M11 from May-September, which leading to the an underestimation of O₃ levels. In EA3, spatially averaged O₃ concentrations often differ by more than 20 ppbv in the individual models. The highest levels of intermodel variability jon O₃ values appeared from May-October, which overestimatinged O₃ levels in comparison torelative to observations by 10-40 ppbv. Interestingly, although M8, M9 and M14 exhibited a -similar magnitudes with observations form June-September, they significantly underestimated observations in other months by 200-300%. A detailed investigation is required in future studies.

3.3 Diurnal concentrations of surface O₃

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Sub-regional O₃ diurnal variations are shown in Fig. 4. In general, model results for three sub-regions exhibited a larger spread with a magnitude of 10-50 ppbv throughout_across the diurnal cycle than that those observed in Europe and North America (Solazzo et al., 2012). The Summer Ense O₃ in summer levels exhibited a systematic pattern of overestimation (20 ppbv) throughout the diurnal cycle in EA1. This indicatesd that the models had difficulty dealing with measuring summer O₃ levels infor the North China Plain. Compared with to summer conditions, there was only a slightly systematic overestimation of Ense O₃ levels in was observed for the other seasons (3-5 ppbv). In EA2, Ense O₃ levels generally agreed with the observations in summer, autumn and winter observations. In particular, the O₃ maximum occurring at around noon was reproduced, reasonably reproduced. There was oQnly a 3-5 ppbv overestimation was observed from during 16:00-23:00 and in early morning (6:00-10:00). In the spring, a systematic overestimation of Ense O₃ walues exited was observed across in the whole diurnal cycle (5-10 ppbv). In EA3, Ense captured the minor small-diurnal variations of of a cross all in-four

seasons, but significantly overestimated observations <u>for the in-</u>summer and autumn (5-20 ppbv). In <u>the</u> spring and winter, differences between Ense and observations <u>fellwere</u> within 5 ppbv.

Among Of all of the models, M11 exhibited the best model performance level in measuring on peak daily O_3 concentrations of 60 ppbv from 14:00-16:00 in EA1, but it still overestimated nighttime O_3

- 5 <u>levels</u> by 10 ppbv. Compared to their with performance in measuring summer patterns, the models² performances had _ performed a significantly better improvement in measuring winter conditions because ofdue to the weak intensity of photochemical reactions, except in case of M2, M10 and M8. Differences between observations and most simulations in for both the nighttime and daytime were fell within 5 ppbv. These differences e contrast of in the models' performances between the summer and winter imply ied
- that the variety of parametrizations on chemistry parametrizations applied to in different models partly explained the intermodel variability of simulated O₃ levels in EA1 (North China Plain). ForIn EA2, the majority of models agreed well with the diurnal variations occurring in the summer and autumn. However, most models exhibited had a tendency to overestimate the O₃ concentrations forin both both the daytime and nighttime in the spring. The overestimated magnitudes exceeded 10 ppbv and 25 ppbv (out-of observed values of 20-35 ppbv) for the in-nighttime and daytime, respectively. M11 reproduced d the observed O₃ levels for in the spring, but underestimated O₃ levels for the in-summer and autumn. ForIn EA3, the significant levels of intermodel variability persisted still existed throughout the year. AThe amplitudes of intermodel variability except for those of M8 and M14 reached approximately 20 ppbv and approximately-10 ppbv in the spring-summer and autumn-winter, respectively. M8 and M14

3.4 Error statistics on surface concentrations

In this section, we present statistics <u>concerning on</u> the <u>models'</u> performance <u>levels of the models</u> based on monthly values. <u>They Values</u> are calculated <u>by with equations shown</u> in Appendix A. On a yearly basis, all models <u>showed observe</u> the highest (0.8-0.9) and lowest (0.1-0.6) correlation coefficients for O₃ <u>forin</u> EA1 and EA2, respectively (Table 2). <u>The hH</u>igh correlations <u>were observed</u> in EA1-were mainly because the summer-maximum and winter-minimum seasonal cycles <u>areis the</u>-typical pattern inof polluted regions that were well represented in all <u>of</u> the participating models. In general, Ense <u>performed</u>-performeda better <u>performance level</u> than the individual models for in representing NO₂ <u>forin</u> East Asia, reproducing the observed seasonal cycles and magnitudes. However, Ense did not always exhibit_ed-a superior performance in measuring for O₃ levels over certain individual models forin East Asia, which <u>standswas</u> in contrast to its performance in <u>for</u> Europe (Table <u>12</u>). M7 and M11 agreed well with observations forin EA1 and EA2, while ENSE tendsed to overestimate O₃ concentrations forin May-

- 5 September in EA1 and <u>for</u> January-September in EA2. Loon et al. (2007) indicated that E<u>nseNSE</u> exhibits ed-a superior performance level only when the spread of ensemble-model values <u>iwas</u> representative of the O₃uncertainty-of O₃. This indicates that most models doid not reflect this uncertainty or missed key processes <u>in-of</u> MICS-Asia III.
- The largeConsiderable overestimation s made byof most of the models forin May-September led to
 high NMB (0.25-1.25) and RMSE (10-33 ppbv) values for in EA1. M11 generatedhed the lowest NMB (0.09) and RMSE (9.46 ppbv) among-values of the examined models. ForIn EA2, M9 and M10 generatedhed strongerlarger correlations than the other models. However, their corresponding NMB and RMSE values were also the highest. These findings is impliedy that systematic model biases are present existed in these two models. M7 exhibited a lower NMB and RMSE values than the other models, but its correlation was measurements as only 0.29. ForIn EA3, the correlations exhibited the largest degree of intermodel variability among across all sub-regions, ranging from -0.13-0.65. M7 generatedshowed the lowest NMB and RMSE and This is likely eaused bydue to the cancelling effect of its overestimation for theim summer and underestimation in for other seasons (Fig. 3).
- For NO, correlations of model correlations forin EA1 ranged from 0.57-0.68, which
 indicatedshowing that all of models did a good job ineffectively reproducinged -the-spatial variability
 inof NO for in-this sub-region (Table 3). The NMBs indicated underestimation by the models except in
 case of M8, which mostly occurred for thein winter. This underestimation can be partly was attributed to
 the coarse model horizontal resolution (45 km) used in the MICS-Asia III, which hardly reproduced
 concentrations of short-lived species (e.g., NO). In contrast to most of the other models, M8
 overestimated NO concentrations in for all three sub-regions. It is noted that NO observations of NO for EA3 were too low (<0.3 ppby) in EA3-to be discussed in this study.

Table 4 shows the statistics of on the models' performance levels in measuring for NO_2 levels. In general, most of the models exhibited a better performanceed better inlevels for representing NO_2 than

 O_3 and NO in-for EA1. The-NMBs ranged from -0.28-0.32, which were much lfalling far below ower thanthose measured for O_3 ((0.48-1.25). The cCorrelations of were 0.54-0.66 were recorded, implying the models' reliable model-performance levels forin reproducing the sepatial and month to-monthly variability of NO₂ forin EA1. Similar those forto O_3 and NO, the correlation coefficients for NO₂ in EA2 remained low._.-Thus, a dedicated investigation of O_3 , NO and NO₂ levels i in EA2 is urgently needed, but falls beyond the scope of this study. In EA3, correlation coefficients ranged from 0.5-0.72. The-NMBs and RMSEs except for those of M8 ranged from -0.42-0.46 and 0.91-1.79 ppby, respectively.

3.5 Vertical profiles of O₃

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Fig. 5 <u>shows</u> the vertical profiles of observed and simulated O₃ <u>levels in for East Asia forin the</u>
summer and winter. Ensemble means (Ense) <u>presented show an</u> underestimations and overestimations
<u>offor EA2 O₃ levels for the in-</u>middle (500-800 hpPa) and lower (below 900 hPpa) troposphere,
respectively. In <u>the winter, the underestimations was even extended to 200 hPpa. MThe magnitudes of</u>
underestimations and overestimations reached 10-40 ppbv and 10-20 ppbv, respectively. <u>ForIn EA3</u>,
Ense reproduced the vertical structure of ozone <u>for in-</u>both <u>the</u> summer and winter. An overestimation <u>of</u>
existed less thanbelow 800 hpahPa in summer, with a magnitude of 10-20 ppbv was observed for the summer.

A large-High levels of intermodel variability of in O₃ exceeding above 300 hPa is evident across all in all-sub-regions, which is attributable to the various-varied different-top boundary conditions among applied by the models. However, this large-considerable variability was not transmitted to the middle troposphere (400-600 hPa), in which O₃ concentrations were consistent among across the models. In the lower troposphere, a small-minor level of intermodel variability of below 900 hP appeared in the winter appeared below 900 hPa-in three sub-regions, and slowly decreased with height. The mMean standard deviations (SD) of models-(G) of below 900 hPpa were recorded as 7.6 ppbv, 6.9 ppbv and 6.0 ppbv in for EA1, EA2 and EA3, respectively, which-coveringed 18.3%, 15.0% and 15.4% of mean O₃

concentrations. In 700-900 hPpa, σ-SD levels decreased to 5.4 ppbv, 4.4 ppbv and 4.8 ppbv in for EA1,
 EA2 and EA3, 12.2%, 9.4% and 10.8% of mean O₃ concentrations, respectively.

In the lower troposphere, the intermodel variability in the summer were was generally higher than those that in the winter. In polluted regions (EA1), σ -SD levels reached 16.3 ppbv (20.8 % of mean

concentrations) in <u>the</u> summer, greatly exceeding those_-in winter (6.2 ppbv, 15.2%). Various vertical structures of O₃ were found below 700 hPa in summer. O₃ concentrations slowly increased with height in M8 and M11, but they mixed well in the PBL and decreased from 800 hPa to 700 hPa in the other models. Akimoto et al. (2019) found that the parameterization on downward O₃ transport from the upper boundary layer contributed <u>a lot toconsiderably to the</u> discrepanciesy between M1, M6 and M11. In EA2, vertical structures of O_{3*} among models were <u>found to be</u> consistent, but concentrations differed more than those in EA1. <u>SDe values</u> covered 22% of mean concentrations.

4. Multi-model ensemble O₃ and comparison with MICS-Asia II

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4.1 Spatial distribution of single model and multi-model ensemble O₃

Fig. 6 shows that the spatial distributions of MICS-Asia III ensemble mean surface O₃ <u>values</u> (Ense) and the coefficient of variation (CV). The CV is defined as the standard deviation of the modeled O₃ divided by the average. The larger the <u>value of CV value</u>, the lower the <u>degree of consistency</u> among the models. In For the summer, ENSE-Ense predicted the an elevated O₃ concentration belt in the middle-latitudes (30°-45°N). A region of O₃ in excess of 60 ppbv stretched across the North China Plain and East
China East-Sea, which was muchfar -higher than values in-MICS-Asia II (45-50 ppbv) <u>values for the year of for 2001</u> (Han et al., 2008). In other seasons, the O₃ <u>values</u> distribution shows higher O₃ over the ocean than in eastern China, reflecting the O₃ titration from high NO_x emissions. Due to the stratospheric injection, surface O₃ over the <u>Tibet Pplateau</u> remained <u>at a-high levels</u> in the wholethroughout the year, ranging from 50 to 65 ppbv. The seasonal cycle of surface O₃ levels determined from in-Ense in-via MICS-Asia III agreeds with that in-observed from MICS-Asia II, but O₃ levels in polluted regions were higher (Han et al., 2008).

The CV ranged from 0.1-0.6 in East Asia. The highest values were found in EA1 in <u>the</u> winter. These high values in <u>the</u> low-latitude western Pacific (10°S-15°N) and Indian Ocean<u>s</u> were likely caused by the treatment of lateral boundaries in <u>the</u> models. <u>In-For</u> MICS-Asia III, M7, M8 and M9 employed the default <u>model</u> configurations<u>of</u> models, and the others employed outputs of <u>the</u> GEOS-Chem/CHASER/MOZART-GOCART global model. Compared <u>with-to those MICof MICS</u>-Asia II, the

CVs <u>in-for the Asian continent except for the winter remained <u>at a-similar levels</u> in this study (0.1-0.3) (Carmichael et al.,2008).</u>

Although all of the models similarly predicted the emergence of an elevated summer O3 concentration belt in the middle-latitudes ($30^{\circ}-45^{\circ}N$), the magnitudes of the enhanced O₃ were levels 5 varied between different among the models (Fig. 7). M5 predicted the highest O₃ concentrations of 60-90 ppbv in-for the North China Plain (EA1) and for its outflow pathways including the -Bohai Sea, East China Sea, Korea, Japan and the Sea of Japan (Locations are shown in Fig. S1 in the supplementssupplementary section), whereas M8 predicted the lowest levels of 35-50 ppbv. Overhangs of 30 ppbv contour lines extending into the nNorthwestern Pacific in-among the Asian continent outflow 10 plume differed considerably among-between the models. The A plume of 30 ppbv or higher O₃ levels was observed from in M1-M6, M13 and M14, reachinged further south and east of Japan (135°E, 20°N), than those of M8, M10 and M11 (120°E, 30°N). In-From MICS-Asia II and HTAP, differences of in the frequency of marine air masses from the western Pacific Ocean were thought to be a possible cause of O₃ discrepancies observed y over oceans among between the models because of due to different 15 meteorological drivers (Han et al., 2008). Forln MICS-Asia III, the winds fields reflected inby the models were similar becausedue to the use of the models the same or similar meteorological fields (Fig. S2 in the supplementary sections). Hence, Ttheseis inconsistenciesy among between the models have resulted from the combined influence effects of a series of factors-, that included ing the diversity in-of condensed gas-chemical mechanism and heterogeneous chemistry. Li et al. (2015) found that the chemical 20 production was to be the dominated controlling factor of O_3 along the outflow pathways near the North China Plain in the summer, rather than lateral and top boundary conditions. Impact of aerosols on ozone in these regions were are frequently reported in Olson et al. (1997) and Li et al. (2018), by to altering photolysis rates and heterogeneous chemistry patterns. DThe detailed comparisons on of parameterizations of these processes in models are needed in future intermodel comparison projects 25 focused on-in Asia.

In <u>the</u> winter, the distribution patterns of O_3 were quite <u>alike similar between the among</u> models, with high concentrations <u>observed</u> over parts of western China, northeastern India and the western Pacific from the East China Sea to south<u>ern-of</u> Japan (Fig. S3 in the supplement<u>ary sections</u>). In spring and autumn (Fig. S4 and Fig. S5 in the supplements), O_3 concentrations were generally higher than <u>they were</u> in <u>the</u> winter <u>acrossin</u> the whole model domain <u>because ofdue to</u> the enhancement of solar radiation or stratosphere-troposphere exchanging fluxes of O_3 . <u>All of the models exhibited an enhancement of O_3 </u> <u>levels over southern Tibet, northeastern India and the western Pacific, generally echoing patterns</u>

5 <u>observed in the winter</u>A major feature consistently produced by all models was the enhancement of O3 over southern Tibet, northeastern India and the western Pacific, which was generally similar to that in winter. <u>Increases in The position of O3 observed enhancement further north of Japan arewas</u> comparable with winter.

4.2 Comparison with MICS-Asia II

10 In-From MICS-Asia II, model evaluation on O_3 were conducted in only on sites in the western Pacific. Fig. 8 presents the simulated and observed surface O_3 levels at these monitoring sites in-derived from the phase II and III of the MICS-Asia project. Note that different models were employed in two phases. In general, most of the models captured the major distributions of O₃ at most sites in both MICS-Asia II and III. ENSE results are showed a good consistent ey form March and December of 2001 and 15 2010. The <u>U</u>underestimations of O_3 levels in March at Japan sites (site 4: Sado-seki, site 5: Oki and site 6: Banryu) in Phase II wereas largely improved remedied in Phase III. However, the surface O_3 at observed in western Japan (site 4: Oki, site 5: Hedo and site 6: Banryu) were severely overestimated in July 2010 by 10-30 ppbv. This overestimation has was not been found in Phase II, in for which the differences with from observations was were valued at approximately 5 ppbv. Rural sites in western 20 Japan were are located in the upwind regions of Japanese domestic emissions, and and are subjected to usually used to capture the impacts of Asian continent outflows. The oOverestimated O_3 ivalues n for North China Plain (EA1) in Phase III contributed considerablya lot to the enhanced concentrations measured for at-sites of western Japanese sites in July 2010. This indicates d that the-transboundary transport from the Asian continent according toin MCIS-Asia III was likely overestimated eompared 25 relative to with that in-measured from MICS-Asia II.

5. Discussions

In <u>reference to MICS-Asia II</u>, Han et al. (2008) <u>guessed hypothesized</u> that the diversity of <u>variations</u> <u>in</u>_meteorological fields, dry deposition, PBL, model treatment of chemistry and other physical processes <u>had</u> contributed to model biases <u>with-in relation to</u> observations and the-intermodel variability.

- 5 Quantifying the contributions of these processes is one effective way tocan help explain model biases by through sensitivity simulations. But However, this task comes with required a tremendous amount of computational costs when applied to for 14 models. TheA qualitative analysis of potential causes of these processed based on by comparisons between of models and observations on these processes is essential to narrow selecting sensitivity simulating scenarios for the next phase of MICS-Asia. In MICS-
- 10 Asia III, common input data (emission and meteorology) provide a good chancebe effectively used in for this qualitative analysthis qualitative analysis based is on model parameterizations. We evaluated the models on dry depositions, PBL and chemistry by collecting their observations (dry deposition velocity and PBLH)-as-much as possible. This work was not conducted in-under MICS-Asia II and is believed intended to be-helpful for-model developers to improve model performance in-for East Asia.

15 **5.1 Dry depositions**

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Previous studies <u>revealed_show</u> that dry deposition processes <u>are-serve as</u> the key net sink of O_3 , accounting for <u>about-roughly</u> 25% of total removed from the troposphere (Lelieveld and Dentener, 2000). The uncertainty of dry deposition in CTMs is still high because many processes are heavily parameterized in models (Hardacre et al.,2015). In this study, the simulated dry deposition velocities of O_3 were compared. Simulated deposition velocities were calculated from Eq. (1):

 $V_d = F/C \tag{1}$

<u>w</u>Where *F* and *C* represent the simulated dry deposition flux and surface O_3 concentrations, respectively. We determined the spatial mean dry deposition velocities at from stations in each sub-region.

Fig. 9 presents the simulated and observed monthly spatial mean dry deposition velocities of O₃.
 In For EA1, ensemble mean values underestimated overestimated observed dry deposition velocities of O₃ (v_d) in for August-September, but still fell into-within the range of the observed standard deviation. This shows that other factors rather than dry deposition could play important roles in overestimations of

August-September O_3 values in EA1. In October-November, simulated v_d apparently underestimated observations by 30-50%. Among the models, This underestimation could contribute to the overestimation of O_3 -concentrations in summer discussed in section 3.2. Tthe lower dry deposition velocities in May-July from-for M1, M2, M4 and M6 than that of M11 partly explained higher summer May-July surface O_3 from those simulations than that from-in M11. However, M13 and M14 still

- <u>May-July</u> surface O₃ from those simulations than that <u>from-in M11</u>. However, M13 and M14 still produced high O₃ concentrations in May-September although their dry deposition velocities were similar to that of M11(Fig. 3). This suggested that there were other factors besides dry deposition playing important roles in the overestimation of summer O₃-in EA1. In October November, simulated v_d apparently overestimated observations by 30 50%. Notably, our observations were made on grassland, which covers ~20% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1. There are few v_d observations on agriculture crops (50% of the land area in EA1.
- the land area) in China. Hardacre et al. (2015) reported O₃ dry deposition measurements on crops in Europe and simulated O₃ dry deposition in 15 global models. Both observations and simulations showed that O₃ dry deposition velocities on agriculture crop class were quite similar to those of grassland, showing uncertainties related to be the representativeness of measurement sites used in this study did not affect our conclusions.

InFor EA2, similar features with as those of EA1 wereare found. M1, M2, M4 and M6 were quite consistent with each other, with a seasonal cycle and a of spring minimum. M11, M12 and M14 showhad no obvious signs of seasonal variability, with a magnitude of 0.1-0.2 cm/s. SThe seasonal patterns in M13 are was considerably different from the those of the other models, exhibiting a maximum in April-September with higher dry deposition velocities (0.5 cm/s). The performance of the models for dry deposition velocities was not always consistent with O₃ concentrations. For example, O₃ concentrations in M13 still-remained high levels under higher dry deposition velocities conditions.

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In EA3, most stations were <u>located in</u> remote oceanic sites, and few dry deposition observations were <u>conductedmade</u>. <u>ThusSo</u>, we collected observations <u>in-from</u> other oceanic sites to evaluate model performance (Helmig et al., 2012). Ense <u>values fofor</u> v_d agree <u>reasonablyd</u> well with with observations <u>reasonably</u> (Fig. 9). Both observations and simulated v_d <u>values</u> showed a July-September maximum with a magnitude of 0.02-0.03 cm/s. Park et al. (2014) revealed found that surface O_3 levels in EA3 were to be more sensitive to dry deposition parameterization schemes in CTMs. O_3 measured from on-oceans differed by 5-15 ppbv in East Asia resulting from due to the use of different various dry deposition parameterization schemes. Thus, more observations are needed over oceans in EA3 to decrease mitigate the uncertainties on O_3 simulations uncertainties.

5.2 Relationships between surface NO_x and O₃

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In general, surface O₃ mainly comes from the photochemistry <u>processes</u> involving NO_x and VOCs in polluted regions. Theoretical and simulation results showed that O₃-production increased almost linearly with the NO_x-increase under NO_x-sensitive conditions and remained relatively unchanged or even decreased in NO_x-saturated (often called "VOCs limited") conditions (Kirchner et al.,2001; Sillman and He et al., 2002; Tang et al., 2010). Recent observations found that regional O₃ in the North China (EA1) and Pearl River Delta (EA2) was changing from NO_x-limited to NO_x-saturated regions (Jin et al., 2015). Examining the O₃-NO_x relationships is <u>a good way of effectual to</u> investigating sources of intermodel variability and model errors concerning on O₃ chemistry in East Asia. Fig. 10 presents the O₃ concentrations as a function of NO_x in May-September based on the monthly daytime (8:00-20:00) mean observed and simulated results at for the stations shown in Fig. 1.

In-For_EA1 (North China Plain), observations clearly revealed show that O₃ concentrations decreased with the <u>an</u> increase in NO_x concentration. O₃ concentrations mostly remained <u>at</u> high levels (40-60 ppbv) when NO_x was less than 20 ppbv. This implied that O₂ was under NO_x saturated conditions in EA1 in May September. The slope and intercept of the regression line between <u>observed</u> O₃ and NO_x were <u>measured as</u> -0.77 ppbv/ppbv and 59.5 ppbv, respectively. Among the models, M11 were results are in relative agreement with observations, reasonably. The slope and intercept (-1.01 ppbv/ppbv, 63.23 ppbv) were reflect the close to observations. The Oother models showed a higher degree of model bias and intermodel variability <u>i</u>on relationships between O₃ and NO_x. Their slopes mostly ranged from -1.25 ppbv/ppbv to -2.13 ppbv/ppbv, <u>amounting to</u> 1.3-2.8 times of observed slope. Their intercepts were 74.9 -121.2 ppbv, much higher thanfar exceeding observations (59.5 ppbv). Akimoto et al. (2019) calculated the net photochemical production of M1, M6 and M11 and found that weak net chemical production in M11 were mostly responsible for low O₃ rather than those in M1 and M6. This finding is consistent with

the low slope in M11. To reduce the impact of O_3 buildup and transport by consuming NO_x , relationships between O_x (NO_2+O_3) with NO_x was compared (Fig. S7 in the supplementary section). Observed O_x increases with the increase of NO_x levels, with coefficient of determination (R^2) of 0.61. Most of the models (except for M8, M11 and M13) failed to reproduced observed positive correlations between O_x

and NO_x, and their R² only ranged from 0.01-0.08. The slope, intercept and R² of M8 and M11 are relative agreement with observations. Akimoto et al. (2019) calculated the net photochemical production of M1, M6 and M11, and found that weak net chemical production in M11 were mostly responsible for low O₃ than M1 and M6. This is consistent with the low slope in M11. Interestingly, M13 maintained a similar O₃-level at all NO_x-levels (Slope: 0.09), which was different from other models and previous theoretical
 results.-

In-For EA2, all models reproduced observed key patterns in which O_x positively correlated with NO_x . For O_3 -NO_x relationships, M1, M2, M4 and M6 reproduced observed O_3 levels under in low NO_x conditions (< 30 ppbv) but failed to capture the low O_3 under high NO_x conditions (30~40 ppbv), accounting for. This explained the overestimations of these models for O_3 in May-September. By contrast, M8 and M11 produced excessively high NO_x values, <u>which</u> resultinged in their underestimations <u>offer</u> O_3 values. In-For M13 and M14, O_3 concentrations were nearly constant in all levels of NO_x. O_3 was positively correlated with NO_x in M9 and M10, which is-stands in contrast to observations. This finding suggests that more attention is needed when policy makers designate the O_3 -regime (VOCs limited or NOx limited regimes) using M9, M10, M13 and M14.

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Stations in EA3 are mostly located over clean oceans or islands. NO_x concentrations were less than 3 ppbv, which indicated showing the that local chemistry appeared towere not be a key factor shaping of O_3 formation. Thus, we did not discuss examine the simulated O_3 - NO_x relationship further in this study.

5.3 Other factors

Previous studies <u>revealed_show</u> that O₃ precursors are mostly constrained within the boundary 25 layer (Quan et al., 2013). <u>The model evaluation on PBLH model evaluation is essential</u> for the interpretation of model biases with observations. Unfortunately, this <u>evaluation evaluation</u> was not applied<u>eonducted</u> in MICS-Asia II. <u>In 2016</u>, <u>Guo et al. (1996) calculated</u> the <u>PBLH using the bulk</u> <u>Richardson number (Ri) method from the radiosonde network of the L-band sounding system of the</u> <u>China Meteorological Administration (Vogelezang and Holtslag,1996). The system provides fine-</u> resolution profiles of temperature, pressure relative humidity, wind speed and direction. In MICS-Asia III, all selected models exhibited the spring-maximum and winter-minimum season cycle <u>iforn</u> EA1 (Fig. S6 in the supplements<u>ary section</u>), <u>which-capturinged</u> the ma<u>injor pattern of climatological patterny</u> of

- 5 PBLH observations (Guo et al.,2016). The Ense on PBLH only overestimated radiosonde measurements by_100-200 m (~10-15%) and . This is likely caused by due to sampling bias the inconsistency of samples between the models and measurements. The simulation recorded as was the mean value of 12 hours (08:00-20:00), while the average of the measurements was calculated based on a 3-hour periods (08:00, 14:00 and 20:00). For In EA2, the observed PBLH did not varyied as much as that for in EA1, and
- 10 differences between seasons were ranged within 100 m. This pattern was captured by the models. As was observed from Similar as EA1, the simulated PBLH forin EA2 was exceeded the measurements by100-200_m-higher than measurements. Few measurements of remote oceanic sites were conducted in East Asia_were collected. ThusSo, we compared simulations with European Centre for Medium-Range Weather Forecasts Reanalysis Data (von Engeln et al., 2013). Both showed a winter-maximum pattern

15 <u>for</u>of PBLH.

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The East Asia monsoon played an important role in seasonal cycle of O₃ in subregions by the longrange transport. Besides local intensive photochemical productions, the O₃ summer maxima in EA1 were also affected by regional transport from Yangtze River Delta under prevailed summer southern monsoon (-20%) (Li et al., 2016). In EA2, a late maximum of O₃ in September November was quite different from EA1 and EA3. This is largely attributed to the long range transport of O₃ and its precursors in the polluted continental air masses from northern China and photochemical formation under dry and sunny weather conditions in autumn (Zheng et al., 2010). In EA3, the seasonal change of O₃ concentrations was characterized by two peaks in spring and autumn. The first and second peak in Mar Apr and May and June were mainly influenced by the inflow from outside of East Asia and chemically produced O₃ by regional emissions, respectively. In the next studies, we will conduct the intermodel comparison on transport fluxes of O₃ between sub regions over East Asia.

6. Summary

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In the<u>Under</u> MICS-Asia III framework, the evaluation and intercomparison of 13 CTMs wasere conducted with using a wide variety of observations covering two Chinese industrialized regions and the western Pacific, using long-term simulations for the year-2010 with a focus. This study has focused on O_{3_n} , NO and NO₂. In particular, surface O_3 levels in China, was evaluated, which wereas absent neglected in the previous model-intercomparison projects, were evaluated. Large-Considerable levels of intermodel variability of in O_3 existed were observed inacross all subregions over of East-Asia in this study, with model concentrations varying by a-factors of 2 to 3 between different models.

- A model ensemble was conducted produced and evaluated. In general, the model ensemble captured 10 the key patterns of monthly and diurnal O_{3-3} NO and NO₂ in the North China Plain and the western Pacific Rim. It failed to capture the observed seasonal cycle of O₃ for thein Pearl River Delta of China. In-For the North China Plain and western Pacific Rrim, the model ensemble severely overestimated surface O_3 levels infor May-September by 10-30 ppby. This overestimation systematically appeared in both daytime and nighttime. Similarly, the model ensemble tendede had a predominate tendency to 15 overestimate the spring daytime and nighttime O₃ concentrations for the in spring in Peral River Delta. Compared to MICS-Asia II, MICS-Asia III was less prone to underestimatingion of surface O₃ in March at for Japanese sites. However, it predicted too enhanced excessively high surface O_3 concentrations forat western Japan in July, which was not the case in-for MICS-Asia II. In term of O₃ soundings, the ensemble model <u>used</u> in this study reproduced the vertical structure in <u>the</u> western Pacific, but overestimated O_3 20 below 800 hpPa in the summer. For the industrialized Pearl River Delta, the ensemble average presented an overestimation for of O_3 levels form the lower troposphere and underestimations in the middle troposphere. This study revealed We find that the ensemble average of 13 models for O3 didoes not always exhibit a superior performance better than to certain individual models in for East Asia in, which contrasted towith its their performance forin Europe. This suggested that the spread of ensemble-25 model values had does not represented all uncertainties inof O_3 levels or that most models in MICS-Asia III models missed key processes. Unlike In contrast to the performance levels for O₃, ENSE demonstrated
 - superior performance level better than individual models for NO2 in East Asia.

MICS-Asia II <u>guessed_outlines_some_potential reasons_causes_</u>of variabilit<u>yies</u> among models. Quantifying the contributions of these processes to O₃ concentrations <u>serves as an is one-</u>effective way to explain model biases <u>throughby</u> sensitivity simulations. <u>However, _But</u>-this<u>would incur</u>-required a tremendous <u>amount of</u>-computational costs<u> for-when applied to</u> 14 models. In this study, we conducted

- 5 a qualitative analysis on of potential causes by comparingson between models and observations for on these processes to <u>identifynarrow</u> sensitivity simulating scenarios for <u>the</u> next phase of MICS-Asia. The <u>Our</u> comparisons show <u>revealed</u> that the ensemble model <u>overunder</u>estimated observed dry deposition velocities of O₃ forin August-September in North China Plain, which showing that other factors rather than dry deposition <u>could</u> may contribute to the overestimation of simulated O₃ concentrations in the
- 10 summer. For theIn western Pacific, simulated v_d values agreed with observations reasonably well. Photochemical treatment in models may contributed to the O₃ overestimations in North China Plain. The studied Mmodels captured the major pattern of climatologically pattern of of PBLH observations form three subregions of ver East Asia. More evaluations of of turbulent kinetic energy in PBL is urgentare needed to for assess the vertical mixing in future studies.

15 Author contribution:

JL, ZW and GC conducted the study design. JL, TN, BG, KY, JF, XW, QF, SI, HL, CK, CL, MZ, ZT, MK, HL contributed to modeling data. ML, JW, JK and QW provided the emission data. LK helped with data processing. HA, GC and ZW were involved in the scientific interpretation and discussion. JL prepared the manuscript with contributions from all co-authors.

20 **Competing interests**:

The authors declare that they have no conflict of interest.

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Appendix A. Statistical Measures

Defining y_{ij} and Obs_{ij} modeled and observed <u>the ith monthly</u> concentrations of air pollutants at the jith

5 station, having mean value \bar{y} and \overline{obs} . m and n represent the numbers of stations and months.

Correlation coefficient (R)

$$R = \frac{\sum_{j=1}^{m} \sum_{i=1}^{n} (y_{ij} - \bar{y})(obs_{ij} - \overline{obs}) \sum_{i=1}^{m} (y_{i} - \bar{y})(obs_{ij} - \overline{obs})}{\sum_{j=1}^{m} \sum_{i=1}^{n} (y_{ij} - \bar{y})^2 \sum_{i=1}^{m} (y_{i} - \bar{y})^2 \sum_{i=1}^{m} ((obs_{ij} - \overline{obs})^2 \sum_{j=1}^{m} \sum_{i=1}^{n} ((obs_{ij} - \overline{obs})^2)}$$
(A1)

Root mean square error (RMSE):

Normalized Mean Bias (NMB)

$$\text{RMSE} = \sqrt{\frac{\sum_{i}^{n} (y_i - Obs_i)^2 \sum_{j=1}^{m} \sum_{i}^{n} (y_i - Obs_{ij})^2}{n}}$$
(A2)

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$$NMB = \frac{\sum_{i=1}^{n} (y_i - Obs_i) \sum_{j=1}^{m} \sum_{i=1}^{n} (y_{ij} - Obs_{ij})}{n \times \bar{y} \times \overline{Obs}}$$
(A3)

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Table and Figure captions:

Table.1 Basic structures, schemes and relevant parameters of the fourteen participating models

25 Table. 2 Statistical analysis for surface O₃ in three subregions over East Asia (R: correlation coefficient; NMB: Normalized Mean Bias; RMSE: Root Mean Square Error)

Table. 3 Statistical analysis for surface NO in three subregions over East Asia (R: correlation coefficient; NMB: Normalized Mean Bias; RMSE: Root Mean Square Error)

Table. 4 Statistical analysis for surface NO₂ in three subregions over East Asia (R: correlation coefficient; NMB: Normalized Mean Bias; RMSE: Root Mean Square Error)

Fig. 1 Model domain of models <u>for</u> except M13 and M14 with <u>the</u> locations of three sub-regions marked in this study. Also show<u>n</u> are <u>the</u> locations of surface monitoring stations <u>used</u> in this study. The meteorological model used <u>for-to</u> provid<u>eing</u> meteorological fields <u>with-for</u> most models also use<u>s</u> this domain. Note that the domains of M13 and M14 are shown in Fig.10.

Fig. 2 Box-plots of observed and simulated annual NO₂ (left column), NO (middle column) and O₃ (right column) frequency distribution by determined from 13 models and, averaged form stations over in EA1, EA2 and EA3, and in time for the whole 2010 year. n represents denotes the numbers of stations. The rectangle represents the inter-quantile range (25^{th} to 75^{th} percentiles). The small star identifies the mean, the continuous horizontal line withininside the rectangle identifies the median, and the whiskers extend

10

15

20

5

between the minimum and maximum values.

Fig. 3 Time series of monthly NO₂, NO and O₃ <u>levels</u> simulated by all models and their ensembles (Ense), in ppbv, averaged over all observed stations <u>acrossin</u> three subregions <u>ofover</u> East Asia (EA1: top row, EA2: middle row, EA3: bottom row). Observations are <u>denoted also shown</u> by the black line. n represents the numbers of stations.

Fig. 4 Seasonal mean diurnal cycle of surface O_3 , in ppbv, as a function of hour<u>s</u>, for all models and their ensembles, averaged <u>acrossover</u> all observed stations in three subregions <u>of over</u> East Asia (EA1: top row, EA2: middle row, EA3: bottom row). Observations are <u>also showndenoted</u> by the black line. n represents the numbers of stations

Fig. 5 Simulated O_3 profiles <u>in-for the</u> summer and winter of 2010, averaged over all observed stations <u>inacross</u> three subregions <u>over-of</u> East Asia (EA1: left column, EA2: middle column, EA3: bottom column). <u>The-O</u>ozonesonde data <u>observe infor</u> 2010 were as taken from the data base stored by World Ozone and Ultraviolet Radiation Data Centre (WOUDC) <u>database</u>

25

Fig. 6 <u>E</u>The ensemble mean seasonal surface O_3 concentrations and CV <u>values</u> for the different seasons. CV is defined as the standard deviation of the modeled fields divided by the average, for the different seasons

Fig. 7 Surface O₃ spatial distribution <u>derived</u> from 13 models for summer 2010 (unit: ppbv).

Fig. 8 The Mmodeled and observed monthly mean concentrations of O₃ at for EANET sites in the phase II (left panel) and III (right panel) of the MICS-ASIA project. The sSolid line represents the ensemble mean. Note that data in MCIS-ASIA II and III data are refer to in the period of March, July and December of 2001 and 2010, respectively. IDs of the mMonitoring sites representsdenote the following: 1: Rishiri (45.12°N, 141.23°E), 2:Ogasawara (27.83°N, 142.22°E), 3: Sado-seki (38.23°N, 138.4°E), 4: Oki

35 (36.28°N, 133.18°E), 5: Hedo (26.85°N,128.25°E), 6: Banryu (34.67°N,131.80°E)

Fig. 9 Simulated and observed monthly O_3 dry deposition velocities (V_d) for M1, M2, M4, M6, M11, M12, M13 and M14 for in three subregions of ver East Asia (EA1: top row, EA3: middle row, EA3: bottom row). TEX, STR, GGSEX and AMMA represents denote observations if reactions of TexAQS06 (7 July– 12 September 2006; north-western Gulf of Mexico), STRATUS06 (9–27 October 2006; the persistent

5 stratus cloud region off <u>the coast of</u> Chile in the eastern Pacific Ocean), GasEx08 (29 February– 11 April 2008; the Southern Ocean), and AMMA08 (27 April–18 May 2008; the southern and northern Atlantic Ocean). Observation<u>al</u> data <u>were takenis</u> from Sorimachi et al. (2003), Pan et al. (2010), and Helmig et al. (2012).

Fig. 10 Scatter plots between <u>for</u> monthly daytime (08:00-20:00) surface NO_x and O₃ <u>forat</u> each station <u>inover</u> EA1 (red), EA2 (green) and EA3 (blue) in May-October, for observations (obs) and models. Also shown are the linear regression equations <u>between for</u> NO_x and O₃ in_EA1 (red) and EA2 (green).

Models	M1	M2	M3	M4	M5	M6	M7	M8	M9	M10	M11	M12	M13	M14
Domain	Ref ^a	Global	10°N -50°N;											
														80 °E -135 °E
Horizontal	45 <u>km</u>	45 <u>km</u>	45 <u>km</u>	45 <u>km</u>	45_km	45 <u>km</u>	45 <u>km</u>	45_km	45 <u>km</u>	45 <u>km</u>	45 <u>km</u>	45 <u>km</u>	$0.5^{\circ} \times 0.667^{\circ}$	45 <u>km</u>
resolution														
Vertical	$40\sigma_p$ levels	$60\sigma_p$ levels	$20\sigma_z$ levels	$40\sigma_p$ levels	$47\sigma_p$ levels	$15\sigma_z$ levels								
resolution														
Depth of first	58_m	58_m	58_m	58_m	58_m	58_m	29_m	58_m	16 <u>m</u>	44_m	48_m	27_m		100_m
layer														
Meteorology	Standard ^b	WRF/NCEP ^b	WRF/NCEP ^b	WRF/NCEP ^b	WRF/	Standard ^b	Standard ^b	GEOS-5	RAMS/NCEP ^b					
										MERRA2 ^b				
Advection	Yamo	Yamo	Yamo	PPM(Colle	PPM	Yamo	5 th order	5 th order	5 th order	5 th order	Walcek	Walcek and	PPM	PPM
	(Yamartino,			lla and			monotonic	monotonic	monotonic	monotonic	and	Aleksic		
	1993)			Woodward							Aleksic	(1998)		
				1984)							(1998)			
Vertical	ACM2	ACM2	ACM2	ACM2	ACM2	ACM2	3 th order	3 th order	YSU	YSU	K-theory	FTCS	Lin and	ACM2
diffusion	(Pleim,2007)						Monotonic	Monotonic				(Forward in	McElroy,	
												Time, Center	(2010)	
												in Space)		
Dry	Wesely	Wesely	Wesely	M3DRY	M3DRY	M3DRY	Wesely	Wesely	Wesely	Wesely	Wesely	Wesely(1989	Wesely	Wesely (1989)
deposition	(1989)	(1989)	(1989)	(Pleim et			(1989)	(1989)	(1989)	(1989)	(1989))and Zhang	(1989)	
				al., 2001)								et al. (2003)		

Wet	Henry's Law	Henry's	Henry's	Henry's	Henry's Law	ACM	Henry's Law	AQCHEM	Easter et al.,	Grell	Henry's	Henry's Law	Henry's Law	Henry's Law
deposition		Law	Law	Law					(2004)		Law			
Gas	SAPRC99(C	SAPRC99	CBM05(Y	SAPRC99	SAPRC99	SAPRC99	RACM-	RACM	RADM2	RADM2	CBMZ	SAPRC99(C	NOx-Ox-HC	SAPRC99
chemistry	arter,2 000)		arwood et				ESRL with	(Goliff et	(Stockwell et		(Zaveri et	arter,2000)	chemistry	
			al.,2005)				KPP	al., 2013)	al., 1990)		al.,1999)		mechanism	
Aqueous	ACM-ae6	ACM-ae6	ACM-ae5	ACM-ae5	ACM-ae5	ACM-ae5	CMAQ	AQCHEM	Walcek and	None	RADM2	Walcek and	-	ACM
chemistry							simplified		Taylor		(Stockwell	Teylor		
							Aqueous		(1986)		et al.,	(1986)		
							chemistry				1990)	Carlton et al.		
												(2007)		
Inorganic	AER06(Bink	AER06	AER05	AER05	AER05	AER05	MADE	MADE	MADE	GOCART	ISORROP	Kajino et al.	ISORROPIAv1.	ISORROPIAv1.7
mechanism	owski and						(Ackermann				IAv1.7(Ne	(2012)	7	
	Roselle,						et al., 1998)				nes et			
	2003)										al.,1998)			
Boundary	GEOS-Chem	Gipson	GEOS-	CHASER	CHASER	CHASER	Liu et al.	CHASER	GEOS-Chem	MOZART	CHASER	CHASER	/	GEOS-Chem
conditions	global model	(1999)	Chem	global	global model	global	(1996)	global model	global model	+	global	global model		global model
	(Martin et		global	model		model				GOCART	model			
	al.,2002)		model	(Sudo et						global				
				al., 2002a,						models ^c				
				2002b)										
Two-way	Off-line	Off-line	Off-line	Off-line	Off-line	Off-line	On-line	On-line	On-line	Off-line	Off-line	On-line	Off-line	Off-line
feedback														

^a Ref represent the referenced domain by MICS-ASIA III project.

^bStandard represents the reference meteorological field provided by MICS-ASIAIII project; WRF/NCEP and WRF/MERRA represents the meteorological field of the participating model itself, which was run by WRF driven by the NCEP and Modern Era Retrospective-analysis for Research and Applications (MERRA) reanalysis dataset.

^cBoundary conditions of M10 are from MOZART and GOCART (Chin et al., 2002; Horowitz et al., 2003), which provided results for gaseous pollutants and aerosols, respectively.

Models	Region	R	NMB	RMSE	Region	R	NMB	RMSE	Region	R	NMB	RMSE
M1		0.89	0.52	19.79		0.48	0.31	14.41		0.57	0.28	15.49
M2		0.90	0.64	18.13		0.10	0.35	15.06		0.66	0.24	13.83
M4		0.87	0.44	18.78		0.41	0.36	14.15		0.01	0.05	17.57
M5		0.87	0.42	19.00		0.30	0.14	13.38		0.34	0.31	19.28
M6		0.90	0.88	25.41		0.15	0.44	17.41		0.52	0.31	16.52
M7	EA1 (n=19) ^a	0.84	0.25	10.03	EA2 (n=13)	0.29	-0.08	11.11	EA3 (n=8)	0.60	0.02	10.97
M8		0.78	-0.47	13.52		0.20	-0.59	19.54		0.55	-0.27	15.32
M9		0.85	0.59	14.84		0.63	0.48	15.69		0.26	-0.09	13.27
M10		0.82	1.24	32.70		0.51	0.72	21.71		0.52	0.11	12.68
M11		0.81	0.09	9.46		0.34	-0.25	13.40		0.65	0.15	12.09
M12		0.89	0.55	18.53		0.36	0.30	13.31		0.57	0.11	11.81

Table 2 Statistical analysis for surface O₃ in three subregions over East Asia (R: correlation coefficient; NMB: Normalized Mean Bias; RMSE: Root Mean Square Error, unit is ppbv)

M13	0.86	0.95	22.69	0.25	0.50	17.04	0.63	0.09	11.04
M14	0.86	0.75	23.33	0.12	0.40	17.01	-0.13	-0.30	20.03
Ensemble Mean	0.89	0.53	15.92	0.38	0.23	11.76	0.52	0.08	11.93
Ensemble Media	0.89	0.56	17.86	0.37	0.31	13.29	0.54	0.11	12.06

a: n represents the numbers of observation stations

Models	Region	R	NMB	RMSE	Region	R	NMB	RMSE	Region	R	NMB	RMSE
M1		0.58	-0.35	20.68		0.22	-0.81	15.16		0.03	-0.35	0.23
M2		0.57	-0.14	23.73		0.14	-0.73	15.21		0.06	-0.27	0.19
M4		0.60	-0.61	22.29		0.18	-0.87	15.72		0.00	-0.39	0.20
M5		0.57	-0.07	20.34		0.24	-0.29	13.80		0.02	0.08	0.35
M6	EA1	0.60	-0.71	23.36	EA2	0.11	-0.89	15.94	$\mathbf{E}\mathbf{A}2$ (m. 9)	0.15	-0.70	0.16
M7	(n=19)	0.63	-0.75	24.91	(n=13)	0.04	-0.78	15.32	EA3 (n=8)	0.27	-0.40	0.15
M8		0.65	0.91	26.89		0.29	1.14	25.06		0.24	3.53	0.94
M9		0.58	-0.82	27.73		0.32	-0.93	16.72		0.22	-0.54	0.14
M10		0.63	-0.90	27.97		0.27	-0.94	16.30		0.39	-0.51	0.14
M11		0.61	-0.34	19.92		0.04	-0.05	14.86		0.41	0.09	0.14

Table 3 Statistical analysis for surface NO in three subregions over East Asia (R: correlation coefficient; NMB: Normalized Mean Bias; RMSE: Root Mean Square Error, unit is ppbv)

M12	0.62	-0.55	21.19	0.13	-0.85	15.64	0.17	-0.48	0.16
M13	-	-	-	-	-	-	-	-	-
M14	0.68	-0.66	22.74	0.01	-0.66	14.77	0.24	-0.50	0.15
Ensemble Mean	0.63	-0.42	20.12	0.21	-0.55	13.58	0.20	-0.03	0.19
Ensemble Media	0.62	-0.58	21.66	0.17	-0.83	15.40	0.17	-0.45	0.16

a: n represents the numbers of observation stations

Models	Region	R	NMB	RMSE	Region	R	NMB	RMSE	Region	R	NMB	RMSE
M1		0.59	-0.18	11.08		0.33	-0.30	12.92		0.54	0.27	1.51
M2		0.64	-0.25	11.30		0.25	-0.43	14.85		0.43	-0.07	1.13
M4		0.65	-0.28	11.62		0.26	-0.32	13.79		0.56	-0.07	1.04
M5		0.57	0.08	10.86		0.30	0.09	12.91		0.60	0.46	1.79
M6	EA1 (n=19)	0.65	-0.22	11.04	EA2 (n=13)	0.23	-0.30	13.86	EA3 (n=8)	0.56	-0.23	0.90
M7		0.59	-0.22	11.42		0.20	-0.25	13.24		0.65	0.19	1.42
M8		0.43	14.32	11.90		0.43	0.15	10.97		0.72	2.38	4.46
M9		0.60	32.30	18.80		0.51	-0.37	12.66		0.49	0.05	1.66
M10		0.61	-10.61	10.65		0.15	-0.08	12.81		0.63	0.06	1.33

Table 4 Statistical analysis for surface NO₂ in three subregions over East Asia (R: correlation coefficient; NMB: Normalized Mean Bias; RMSE: Root Mean Square Error, unit is ppbv)

M11	0.54	0.00	10.82	0.24	0.13	13.56	0.69	0.36	1.58
M12	0.63	-0.16	10.76	0.25	-0.24	13.78	0.61	-0.05	0.91
M13	-	-	-	-	-	-	-	-	-
M14	0.66	-0.12	10.00	0.08	-0.22	14.50	0.60	0.42	0.91
Ensemble Mean	0.65	-0.09	9.89	0.29	-0.18	12.16	0.64	0.25	1.33
Ensemble Media	0.65	-0.13	10.07	0.27	-0.23	12.85	0.59	0.06	1.23

a: n represents the numbers of observation stations

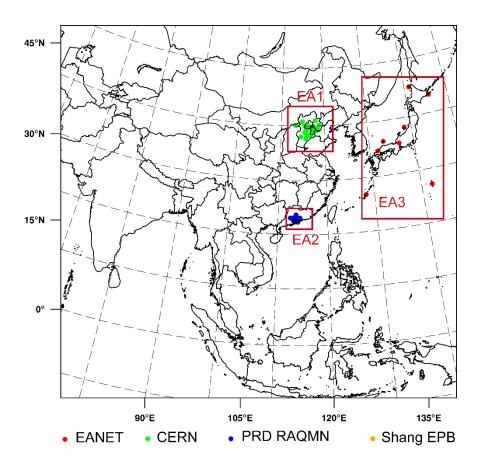


Fig.1 Li et al., 2018

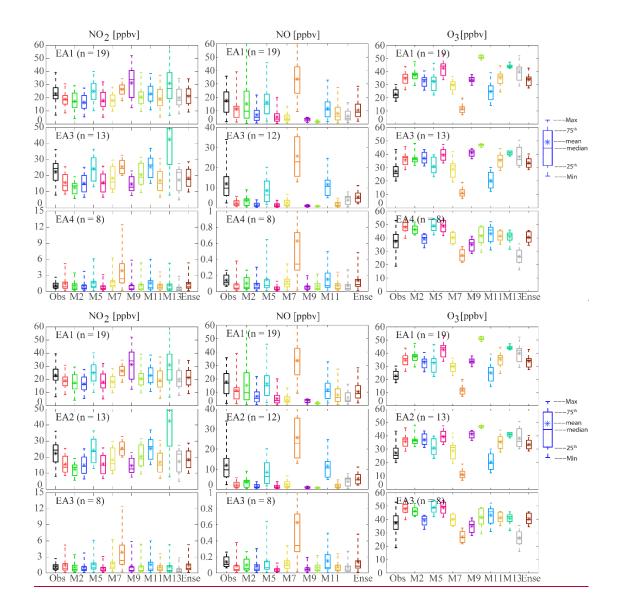
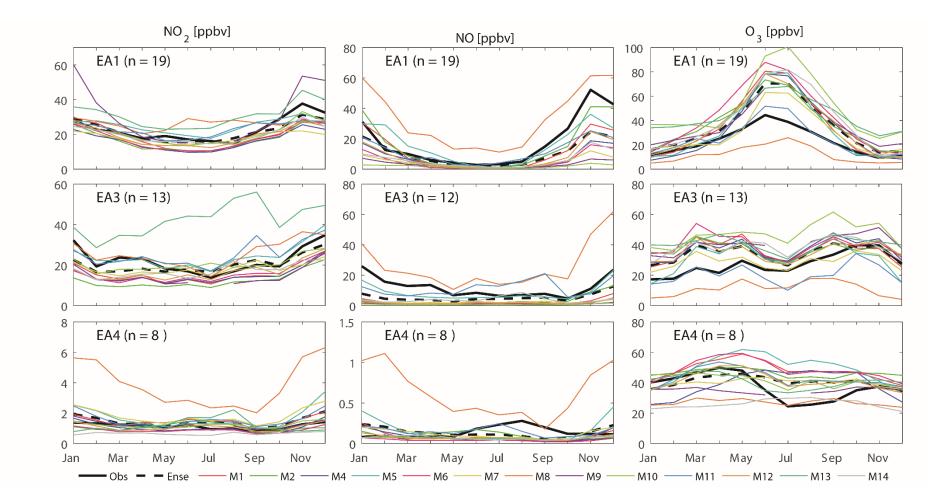


Fig.2 Li et al., 2018



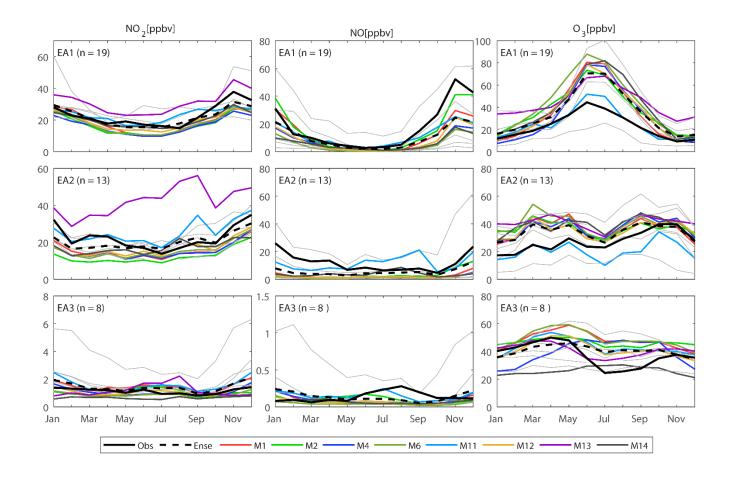
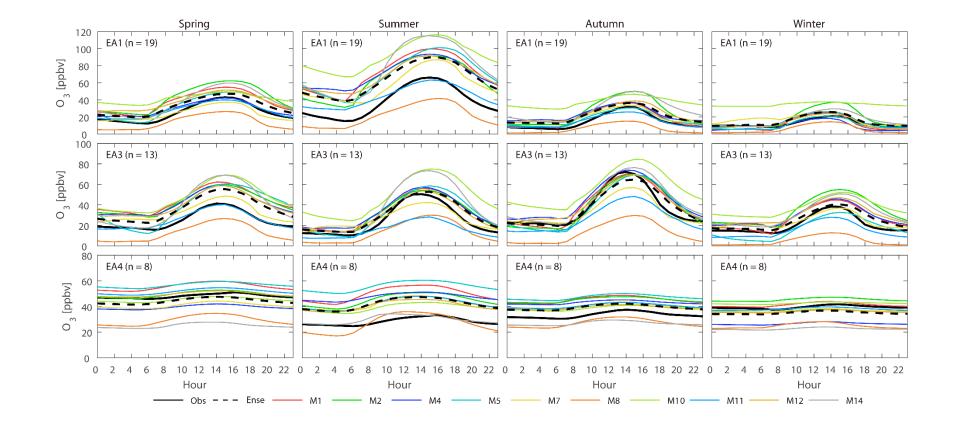


Fig.3 Li et al., 2018



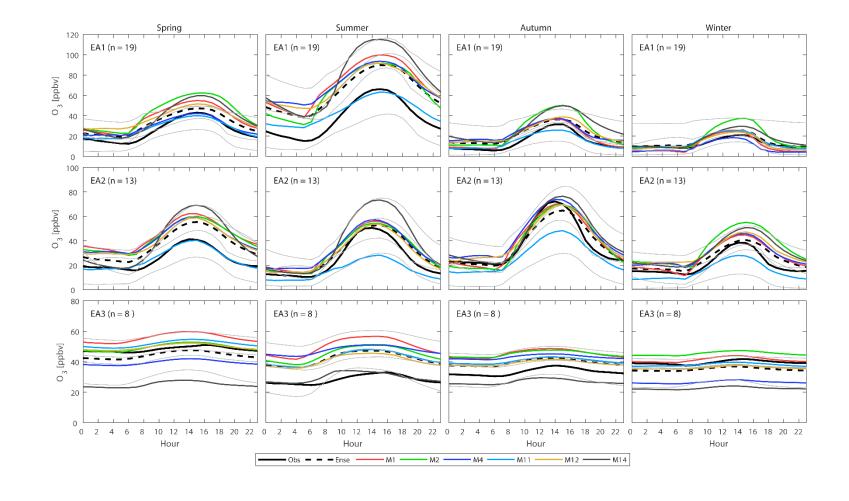
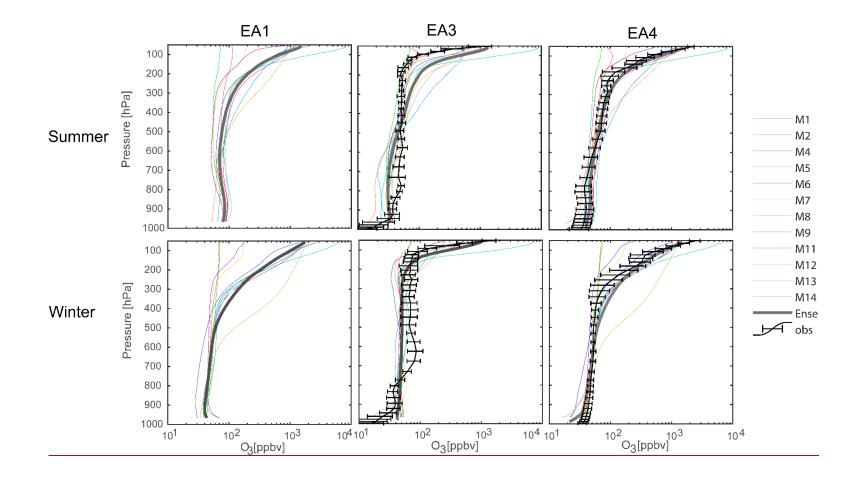


Fig.4 Li et al., 2018



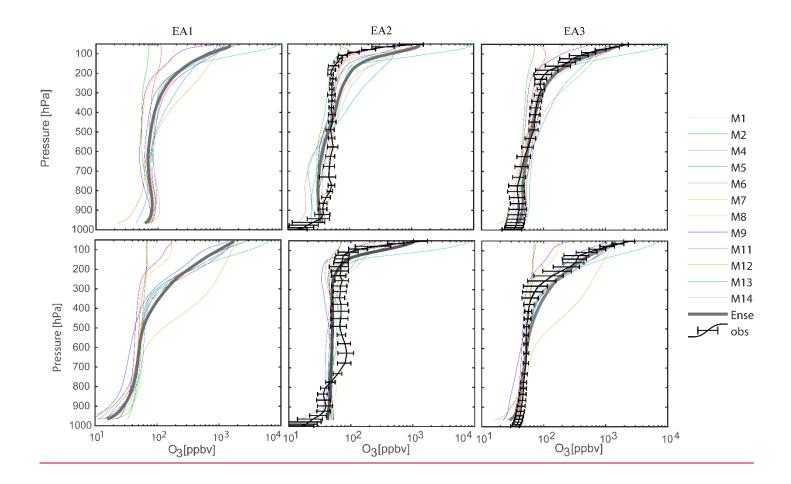


Fig.5 Li et al., 2018

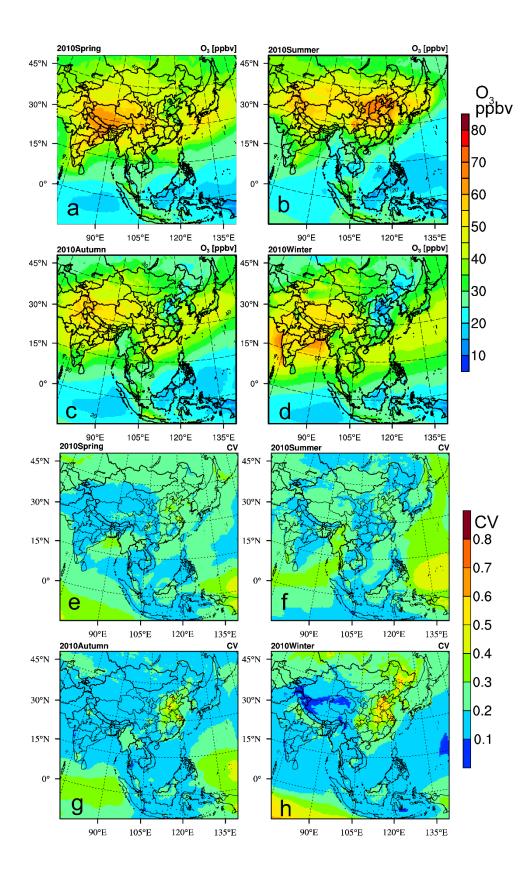


Fig.6 Li et al., 2018

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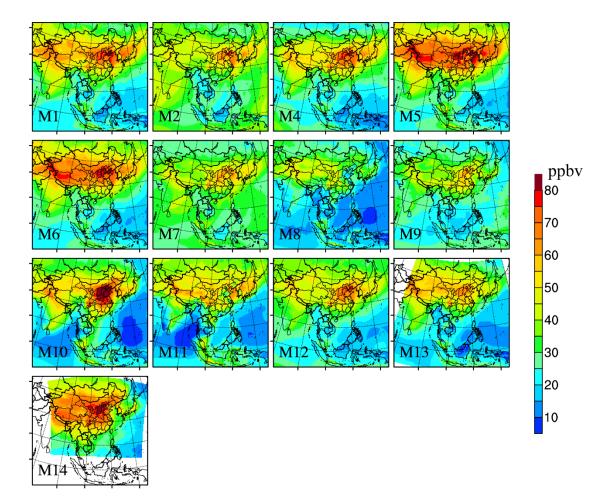


Fig.7 Li et al., 2018

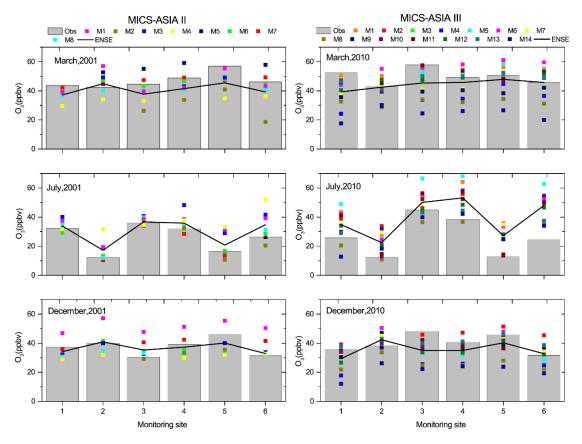


Fig.8 Li et al., 2018

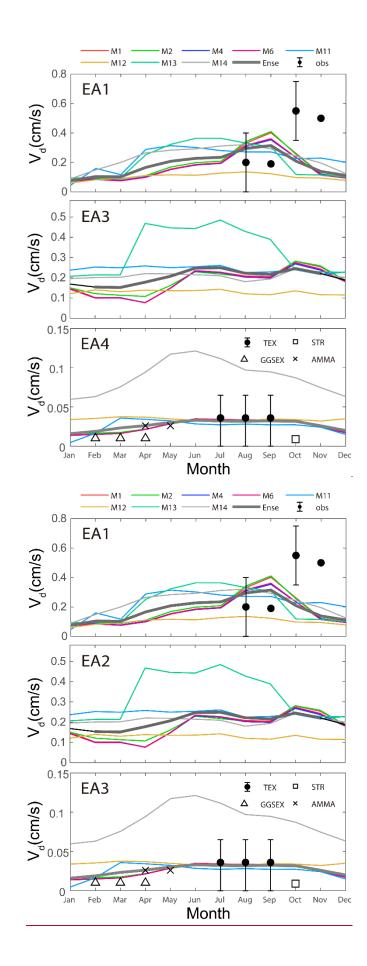
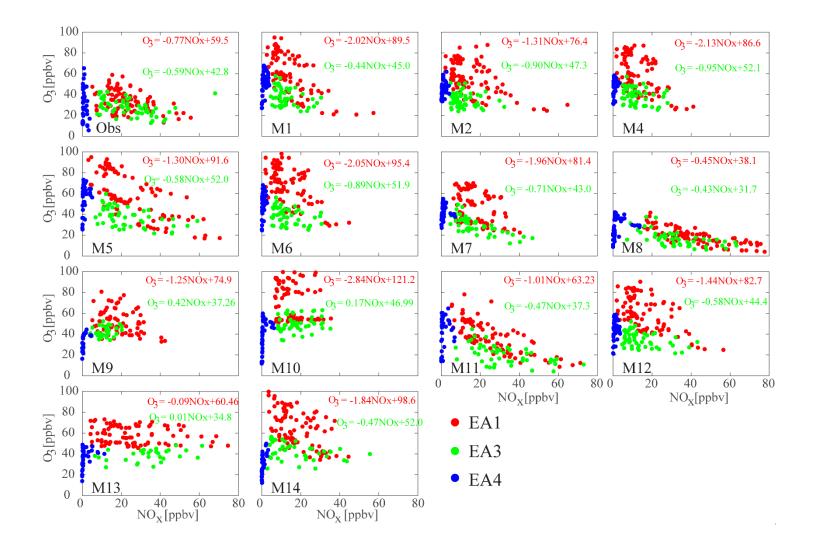


Fig.9 Li et al., 2018



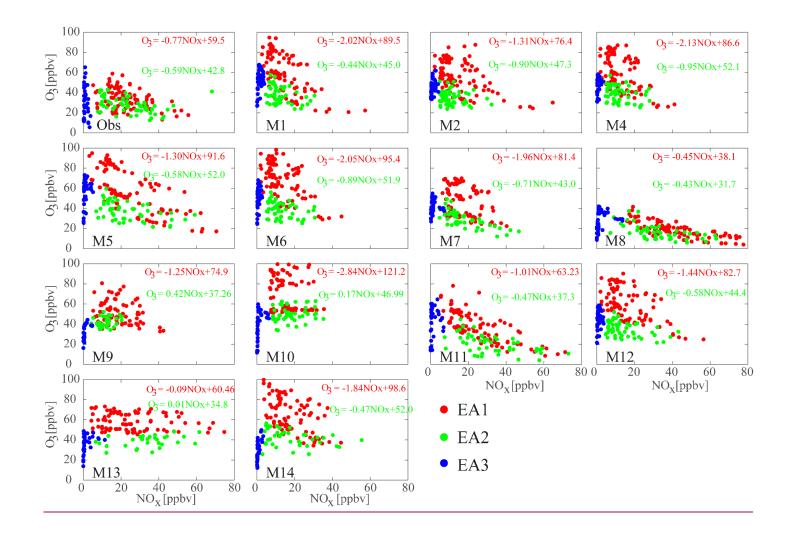


Fig.10 Li et al., 2018