

1 **Modeling the impact of heterogeneous reactions of chlorine on**
2 **summertime nitrate formation in Beijing, China**

3 Xionghui Qiu^{1,2}, Qi Ying^{3*}, Shuxiao Wang^{1,2*}, Lei Duan^{1,2}, Jian Zhao⁴, Jia
4 Xing^{1,2}, Dian Ding^{1,2}, Yele Sun⁴, Baoxian Liu⁵, Aijun Shi⁶, Xiao Yan⁶,
5 Qingcheng Xu^{1,2}, Jiming Hao^{1,2}

6 ¹ State Key Joint Laboratory of Environmental Simulation and Pollution Control,
7 School of Environment, Tsinghua University. Beijing 100084, China.

8 ² State Environmental Protection Key Laboratory of Sources and Control of Air
9 Pollution Complex, Beijing 100084, China

10 ³ Zachry Department of Civil Engineering, Texas A&M University, College Station,
11 Texas 77843-3138, United States

12 ⁴ State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric
13 Chemistry, Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing
14 100029, China

15 ⁵ Beijing Environmental Monitoring Center, Beijing 100048, China

16 ⁶ Beijing Municipal Research Institute of Environmental Protection, Beijing 100037,
17 China

18 *Corresponding author: shxwang@tsinghua.edu.cn_& qying@civil.tamu.edu

19
20 Abstract:

21 A comprehensive chlorine heterogeneous chemistry is incorporated into the
22 Community Multiscale Air Quality (CMAQ) model to evaluate the impact of chlorine-
23 related heterogeneous reaction on diurnal and nocturnal nitrate formation and quantify
24 the nitrate formation from gas-to-particle partitioning of HNO₃ and from different
25 heterogeneous pathways. The results show that these heterogeneous reactions increase
26 the atmospheric Cl₂ and ClNO₂ level (~100%), which further affect the nitrate
27 formation. Sensitivity analyses of uptake coefficients show that the empirical uptake
28 coefficient for the O₃ heterogeneous reaction with chlorinated particles may lead to the
29 large uncertainties in the predicted Cl₂ and nitrate concentrations. The N₂O₅ uptake

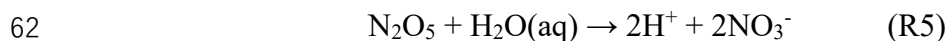
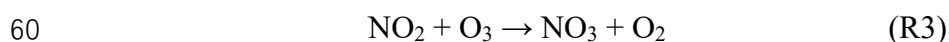
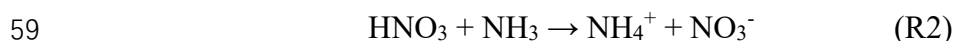
30 coefficient with particulate Cl⁻ concentration dependence performs better to capture the
31 concentration of ClNO₂ and nocturnal nitrate concentration. The reaction of OH and
32 NO₂ in daytime increases the nitrate by ~15% when the heterogeneous chlorine
33 chemistry is incorporated, resulting in more nitrate formation from HNO₃ gas-to-
34 particle partitioning. By contrast, the contribution of the heterogeneous reaction of
35 N₂O₅ to nitrate concentrations decreases by about 27% in the nighttime when its
36 reactions with chlorinated particles are considered. However, the generated gas-phase
37 ClNO₂ from the heterogeneous reaction of N₂O₅ and chlorine-containing particles
38 further reacts with the particle surface to increase the nitrate by 6%. In general, this
39 study highlights the potential of significant underestimation of daytime and
40 overestimation of nighttime nitrate concentrations for chemical transport models
41 without proper chlorine chemistry in the gas and particle phases.

42

43 **1. Introduction**

44 In recent years, nitrate has become the primary component of PM_{2.5} (particulate matter
45 with aerodynamic diameter less than 2.5μm) in Beijing with sustained and rapid
46 reduction of SO₂ and primary particulate matter emissions (Ma et al., 2018; Li et al.,
47 2018; Wen et al., 2018). Observations showed that the relative contributions of
48 secondary nitrate in PM_{2.5} could reach up to approximately 50% during some severe
49 haze pollution days (Li et al., 2018). The mechanism of secondary nitrate formation can
50 be summarized as two major pathways: (1) Gas-to-particle partitioning of HNO₃, which
51 happens mostly in the daytime. The reaction of OH with NO₂ produces gaseous HNO₃,
52 which subsequently partition into the particle phase. The existence of NH₃ or basic
53 particles enhances this process by NH₃-NH₄⁺ gas-particle equilibrium (Kleeman et al.,
54 2005; Seinfeld and Pandis, 2006); (2) Hydrolysis of N₂O₅, which is more important at
55 nighttime. N₂O₅ forms from the reactions of NO₂, O₃ and NO₃ and hydrolyzes to
56 produce particulate nitrate. They can be summarized as reactions R1-R5 (Brown and
57 Stutz 2012):

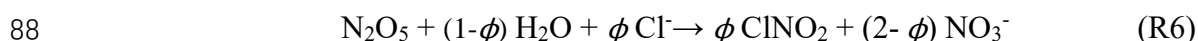




63 In addition to reactions R1 and R5, gas phase reactions of NO_3 with HO_2 and VOCs,
 64 N_2O_5 with water vapor (Tuazon et al., 1983) and the heterogeneous reaction of NO_2
 65 with water-containing particle (Goodman et al., 1999) produce HNO_3 or nitrate as well.
 66 These reactions are listed in Table 2 as reactions R8, R9 and R10.

67 However, chemistry transport models (CTMs) incorporated with these mechanisms
 68 still can't accurately capture the spatiotemporal distributions of nitrate in some studies
 69 in polluted northern China. For example, Chang et al. (2018) showed that the simulated
 70 nitrate concentrations derived from the default CMAQ (version 5.0.2) were
 71 significantly higher than the observations in summer at two sites adjacent to Beijing.
 72 Fu et al. (2016) also found that default CMAQ (version 5.0.1) overestimated the
 73 simulated nitrate concentrations in the Beijing-Tianjin-Hebei region.

74 In recent fields studies, it was found that high particulate chlorine emissions might
 75 have a significant impact on the oxidation capacity of the urban atmosphere and thus
 76 could affect nitrate concentrations. According to the field measurements in June 2017
 77 in Beijing (Zhou et al., 2018), the 2-min averaged concentrations of reactive molecular
 78 chlorine (Cl_2) and nitryl chloride (ClNO_2) reached up to 1000 pptv and 1200 pptv,
 79 respectively, during some severe air pollution periods in summer. The Cl_2
 80 concentrations were significantly higher than those observed in North American coastal
 81 cities affected by onshore flow and the lower atmosphere in the remote Arctic region
 82 (Spicer et al., 1998; Glasow et al., 2010; Liu et al., 2017). During these pollution events,
 83 the corresponding concentrations of N_2O_5 (2-min average) and nitrate (5-min average)
 84 rose from 40 pptv and $1 \mu\text{g m}^{-3}$ to 700 pptv and $5 \mu\text{g m}^{-3}$. To explain the high levels of
 85 ClNO_2 , some studies suggested that reaction R5 should be revised to account for ClNO_2
 86 production from the heterogeneous reaction of N_2O_5 on chloride-containing particles
 87 (CPS) (Osthoff et al., 2008; Thornton et al., 2010), as shown in reaction R6:



89 where ϕ represents the molar yield of ClNO₂. By incorporating this reaction into WRF-
90 Chem, Li et al. (2016) found that the improved model performed better to match the
91 observed nitrate concentrations in Hongkong during 15 November and 5 December
92 2013. However, ClNO₂ could affect the formation of nitrate indirectly by increasing the
93 atmospheric OH after a series of chemical reactions, which are briefly summarized into
94 three steps: (1) the photolysis of ClNO₂ produces chlorine radicals (Cl[•]); (2) the reaction
95 of Cl[•] with VOCs produces peroxy radical (HO₂ and RO₂); and (3) the increased HO₂
96 and RO₂ prompt the formation of OH through HO_x cycle and lead to increased HNO₃
97 production (Young et al., 2014; Jobson et al., 1994). The overall impact of R6 on nitrate
98 remains to be investigated.

99 Another related but unresolved issue is the sources of the high concentrations Cl₂,
100 which could not be explained by the N₂O₅ heterogeneous reaction with Cl⁻ and the
101 subsequent reactions of ClNO₂ in the gas phase. It has been reported that the reactions
102 of gaseous O₃, OH, HO₂, ClNO₂, hypochlorous acid (HOCl), chlorine nitrate (ClONO₂)
103 with CPS can produce Cl₂, which can subsequently photolyze to produce Cl[•]
104 (Knipping et al., 2000; George et al., 2010; Pratte et al., 2006; Deiber et al., 2004; Faxon
105 et al., 2015). However, these heterogeneous reactions on CPS are generally missing in
106 most of the current CTMs and it is unclear whether these reactions will be able to
107 explain the observed Cl₂ concentrations and the overall impact of these reactions on
108 nitrate.

109 Previously, biomass burning, coal combustion, and waste incineration were
110 identified as the main sources of gaseous and particulate chlorine compounds in China
111 from International Global Atmospheric Chemistry Program's Global Emissions
112 Inventory Activity (GEIA) based on the year 1990 and a localized study by Fu et al.
113 based on the year 2014 (Keene et al., 1999; Fu et al., 2018). However, recent source
114 apportionment results of PM_{2.5} in Beijing showed that the contribution of coal
115 combustion had extremely decreased from 22.4% in 2014 to 3% in 2017 with the
116 replacement of natural gas (obtained from official website of Beijing Municipal Bureau
117 of Statistics, available at <http://edu.bjstats.gov.cn/>). Another important source—cooking
118 has received attention as its increasing contribution to PM_{2.5} (accounting for 33% of the

119 residential sector; obtained from the official source apportionment analysis of PM_{2.5} in
120 Beijing in 2017; see <http://www.bjepb.gov.cn/bjhr-b/index/index.html>). Moreover, the
121 high content of particulate sodium chloride was measured from the source
122 characterization studies of PM_{2.5} released from the cooking activities (Zhang et al.,
123 2016). Thus, it is necessary to compile an updated emission inventory for Beijing to
124 include the emissions from cooking and other sources (coal burning, solid waste
125 incineration, biomass burning, etc.) in order to explore the emissions of the chlorine
126 species on atmospheric nitrate formation.

127 In this study, a CMAQ model with an improved chlorine heterogeneous chemistry
128 is applied to simulate summer nitrate concentration in Beijing. Sensitivity simulations
129 are conducted to evaluate the contributions of HNO₃ gas-to-particle partitioning and
130 heterogenous production to aerosol nitrate. The results of this work can improve our
131 understandings on nitrate formation and provide useful information on nitrate pollution
132 control strategies in Beijing.

133

134 **2. Emissions, chemical reactions and model description**

135 2.1 Emissions

136 Generally, the conventional emission inventories of air pollutants in China only include
137 the common chemical species, such as SO₂, NO_x, VOCs, PM_{2.5}, PM₁₀, NH₃, BC, and
138 OC (Wang et al., 2014). Chloride compound emissions were not included. However,
139 the emissions of chlorine species are vital for studying the chlorine chemical
140 mechanism. Recently, the inorganic hydrogen chloride (HCl) and fine particulate
141 chloride (PCl) emission inventories for the sectors of coal combustion, biomass burning,
142 and waste incineration were developed for the year of 2014 (Qiu et al., 2016, Fu et al.,
143 2018, Liu et al., 2018). However, the gaseous chlorine emission was not estimated in
144 these studies. In addition, these studies did not account for the rapid decrease of coal
145 consumption in recent years in Beijing, from 2000 Mt in 2014 to 490 Mt in 2017. More
146 importantly, the cooking source, as one of the major contributors to particulate chlorine
147 in Beijing, is not included in current chlorine emission inventories. Thus, a new

148 emission inventory of reactive chlorine species, which includes HCl, Cl₂ and PCl, were
149 developed in this study for the year of 2017.

150 The emission factor method (equation (1)) is applied to calculate the emissions of
151 these reactive chlorine species from coal combustion, biomass burning, municipal solid
152 waste incineration and industrial processes:

$$153 \quad E_{i,j} = A_i \times EF_{i,j} \quad (1)$$

154 where $E_{i,j}$ represents the emission factor of pollutant j in sector i ; A represents the
155 activity data; EF represents the emission factor. EF for PCl is estimated by $EF_{i,PCl} =$
156 $EF_{i,PM_{2.5}} \times f_{Cl,i}$, where $f_{Cl,i}$ represents the mass fraction of PCl in primary PM_{2.5}.
157 Activity data are obtained from the Beijing Municipal Bureau of Statistics (available at
158 <http://tjj.beijing.gov.cn/>). The Cl₂ emission factor for coal combustion is calculated
159 based on the content of Cl in coal, which had been measured by Deng et al (2017). The
160 PM_{2.5} emission factors and mass fractions of PCl in PM_{2.5} to calculate the emissions of
161 Cl had been described in detail by Fu et al. (2018). PCl in PM_{2.5} for coal combustion
162 and biomass burning are taken as 1% and 9.0%, respectively, based on local
163 measurements in Beijing.

164 Emissions of PCl from cooking, including contributions from commercial and
165 household cooking, are estimated using equation (2):

$$166 \quad E_{PCl} = [N_f \times V_f \times H_f \times EF_{f,PCl} + V_c \times H_c \times N_c \times n \times EF_{c,PCl} \times (1 - \eta)] \times 365 \quad (2)$$

167 where N_f is the number of households, V_f is the volume of exhaust gas from a household
168 stove (2000 m³ h⁻¹); H_f is the cooking time for a family (0.5 h day⁻¹); $EF_{f,PCl}$ and
169 $EF_{c,PCl}$ are the emission factors (kg m⁻³) of PCl for household and commercial cooking,
170 respectively; H_c is the cooking time in a commercial cooking facility (6 h day⁻¹); N_c is
171 the number of restaurants, schools and government departments. V_c is the volume of
172 exhaust gas from a commercial cooking stove (8000 m³ h⁻¹); n is the number of stoves
173 for each unit, which equals to 6 for a restaurant and is calculated as one stove per 150
174 students for each school. η is the removal efficiency of fume scrubbers (30%). $EF_{c,PCl}$
175 is the emission factor (kg m⁻³) of PCl in commercial cooking. These constants are all
176 based on Wu et al. (2018). The PCl fraction in PM_{2.5} from cooking is take as 10%, based
177 on local measurements. HCl and Cl₂ emissions from cooking are not considered in this

178 study.

179 The sectoral emissions of HCl, Cl₂ and PCl are summarized in Table 1. The
180 estimated HCl, Cl₂ and PCl emissions in Beijing are 1.89 Gg, 0.07Gg and 0.63Gg
181 respectively. The Cl emissions estimated for 2014 by Fu et al. (2018) were used for
182 other areas. This simplification is a good approximation because replacing coal with
183 natural gas only occurred in Beijing, and reduction of coal consumption in surrounding
184 regions was generally less than 15%. In addition, strict control measures for biomass
185 burning, cooking and municipal solid waste incineration have not been implemented in
186 most regions yet. Emissions of conventional species for this study period are developed
187 in a separate study that is currently under review and are summarized in Table S1.

188

189 2.2 Chlorine-related heterogeneous reactions

190 The heterogeneous reactions in original CMAQ (version 5.0.1) are not related to
191 chlorine species. In this study, the original heterogeneous reactions of N₂O₅ and NO₂
192 (R5 and R10 in Table 2) are replaced with a revised version which includes production
193 of ClNO₂ from CPS (R6 and R11 in Table 2). In reaction R6, the molar yield of ClNO₂
194 (ϕ_{ClNO_2}) is represented as equation (3) (Bertram and Thornton, 2009):

$$\phi_{ClNO_2} = \left(1 + \frac{[H_2O]}{483 \times [Cl^-]}\right)^{-1} \quad (3)$$

195 where [H₂O] and [Cl⁻] are the molarities of liquid water and chloride (mol m⁻³).

196 In addition, laboratory observations confirmed that the heterogeneous uptake of
197 some oxidants (such as O₃ and OH) and reactive chlorine species (such as ClNO₂, HOCl,
198 and ClONO₂) could also occur on CPS to produce Cl₂. These reactions are implemented
199 in the model and summarized in Table 2 as R13-R18. Note that the products from the
200 heterogeneous uptake of ClNO₂ on CPS vary with particle acidity (Riedel et al., 2012;
201 Rossi, 2003). It generates Cl₂ under the condition of pH lower than 2 but produces
202 nitrate and chloride under higher pH conditions. The reaction rates of the heterogeneous
203 reactions are parameterized as first-order reactions, with the rate of change of gas phase
204 species concentrations determined by equations (4) (Ying et al., 2015):

$$\frac{dC}{dt} = -\frac{1}{4}(v\gamma A)C = -k^1 C \quad (4)$$

205 where C represents the concentration of species; v represents the thermal velocity of
 206 the gas molecules (m s^{-1}); A is the CMAQ-predicted wet aerosol surface area
 207 concentration ($\text{m}^2 \text{m}^{-3}$); γ represents the uptake coefficient. For all gas phases species
 208 (except ClNO_2) involved in the heterogeneous reactions (R6 and R11-R18), a simple
 209 analytical solution can be used to update their concentrations from time t_0 to $t_0+\Delta t$:
 210 $[C]_{t_0+\Delta t}=[C]_{t_0} \exp(-k^1\Delta t)$, where Δt is the operator-splitting time step for heterogeneous
 211 reactions.

212 The rate of change of ClNO_2 includes both removal and production terms, as shown
 213 in equation (5):

$$\begin{aligned} \frac{d[\text{ClNO}_2]}{dt} &= -k_1^1[\text{ClNO}_2] + k_6^1\phi_{\text{ClNO}_2}[\text{N}_2\text{O}_5] \\ &= -k_1^1[\text{ClNO}_2] + k_6^1\phi_{\text{ClNO}_2}[\text{N}_2\text{O}_5]_{t_0} \exp(-k_6^1 t) \end{aligned} \quad (5)$$

214 Assuming ϕ_{ClNO_2} is a constant, an analytical solution for equation (5) can be obtained,
 215 as shown in equation (6):

$$\begin{aligned} [\text{ClNO}_2]_{t_0+\Delta t} &= [\text{ClNO}_2]_{t_0} \exp(-k_1^1\Delta t) \\ &+ \frac{k_6^1\phi_{\text{ClNO}_2}[\text{N}_2\text{O}_5]_{t_0}}{k_1^1 - k_6^1} [\exp(-k_6^1\Delta t) - \exp(-k_1^1\Delta t)] \end{aligned} \quad (6)$$

216 where k_i^1 represents the pseudo first-order rate coefficient of either reaction R17 or
 217 R18, depending on pH.

218 The uptake coefficients γ of gaseous species are obtained from published
 219 laboratorial studies. In the original CMAQ, the uptake coefficient of N_2O_5 is determined
 220 as a function of the concentrations of $(\text{NH}_4)_2\text{SO}_4$, NH_4HSO_4 and NH_4NO_3 (Davis et al.,
 221 2008). In this study, the PCL and NO_3^- dependent parameterization by Bertram and
 222 Thornton (2009) (equation (7)) is used:

$$\gamma_{\text{N}_2\text{O}_5} = \begin{cases} 0.02, & \text{for frozen aerosols} \\ 3.2 \times 10^{-8} K_f \left[1 - \left(1 + \frac{6 \times 10^{-2} [\text{H}_2\text{O}]}{[\text{NO}_3^-]} + \frac{29 [\text{Cl}^-]}{[\text{NO}_3^-]} \right)^{-1} \right] & \end{cases} \quad (7)$$

223 In the above equation, K_f is parameterized function based on molarity of water: $K_f =$
 224 $1.15 \times 10^6 (1 - e^{-0.13[\text{H}_2\text{O}]})$. NO_3^- and Cl^- concentrations are also in molarity. The
 225 uptake coefficient of OH is expressed in equation (8) as a function of the concentration
 226 of PCL following the IUPAC (International Union of Pure and Applied Chemistry,

227 available at <http://iupac.poleether.fr/htdocs/datasheets/pdf/O->
228 H_halide_solutions_VI.A2.1.pdf).

$$\gamma_{\text{OH}} = \min\left(0.04 \times \frac{[\text{Cl}^-]}{1000 \times M}, 1\right) \quad (8)$$

229 where M represents the volume of liquid water in aerosol volume ($\text{m}^3 \text{m}^{-3}$). For frozen
230 particles, the uptake coefficient is limited to 0.02, as used in the original CMAQ model.

231 The uptake coefficients of O_3 , NO_3 , NO_2 , HOCl , ClNO_2 , and ClONO_2 are treated
232 as constants. Among of them, the γ values of NO_3 , NO_2 , HOCl and ClONO_2 are set as
233 3×10^{-3} , 1×10^{-4} , 1.09×10^{-3} and 0.16 based on laboratory measurements (Rudich et al.,
234 1996; Abbatt et al., 1998; Pratte et al., 2006; Gebel et al., 2001). A preliminary value
235 of 10^{-3} in the daytime and 10^{-5} at nighttime is chosen for the O_3 uptake coefficient. The
236 daytime γ_{O_3} is based on the analysis of Cl_2 production rate in a hypothesized
237 geochemical cycle of reactive inorganic chlorine in the marine boundary layer by Keene
238 et al. (1990). The lower nighttime value was also recommended by Keene et al. (1990)
239 who noted that Cl_2 production in the marine boundary layer are lower at night. The
240 uptake coefficient of ClNO_2 depends on the particle acidity, with the value of 2.65×10^{-6}
241 for reaction R17 and 6×10^{-3} for reaction R18 (Robert et al., 2008).

242

243 2.4 CMAQ model configuration

244 These heterogeneous reactions of chlorine are incorporated into a revised CMAQ based
245 on the CMAQ version 5.0.1 to simulate the distribution of nitrate concentration in
246 Beijing from 11 to 15 June 2017. The revised CMAQ model without heterogeneous
247 reactions of chlorine has been described in detail by Ying et al. (2015) and Hu et al.
248 (2016, 2017). In summary, the gas phase chemical mechanism in the revised CMAQ
249 model is based on the SAPRC-11 (Cater et al., 2012) with a comprehensive inorganic
250 chlorine chemistry. Reactions of Cl radical with several major VOCs, which lead to
251 production of HCl , are also included. The aerosol module is based on AERO6 with an
252 updated treatment of NO_2 and SO_2 heterogeneous reaction and formation of secondary
253 organic aerosol from isoprene epoxides. Three-level nested domains with the
254 resolutions of 36km, 12km, and 4km using Lambert Conformal Conic projection

255 (173×136, 135×228 and 60×66 grid cells) are chosen in this work (the domains see
256 Figure S1). The two true latitudes are set to 25°N and 40°N and the origin of the domain
257 is set at 34°N, 110°E. The left-bottom coordinates of the outmost domain are positioned
258 at $x = -3114$ km, $y = -2448$ km. The BASE case (heterogeneous reactions of Cl turned
259 off) and HET case (all heterogeneous reactions enabled) are compared to evaluate the
260 impact of heterogeneous chlorine chemistry on nitrate formation.

261

262 **3. Results**

263 3.1 Model performance evaluation

264 Predicted O₃, NO₂ and PM_{2.5} concentrations from the BASE case simulation are
265 evaluated against monitoring data at 12 sites in Beijing (Table S2) in 11 to 15, June
266 2017. The average NMB/NME values for O₃, NO₂ and PM_{2.5} across the 12 sites are -
267 8%/29%, -7%/59% and -8%/53%, respectively. Predicted hourly Cl₂, ClNO₂ and N₂O₅
268 concentrations were compared with observations measured at the Institute of
269 Atmospheric Physics (IAP), Chinese Academy of Sciences (39.98°N, 116.37°E) using
270 a high-resolution time-of-flight chemical ionization mass spectrometer (CIMS) from 11
271 to 15 June 2017 (for site description, instrument introduction, and analytical method,
272 please refer to the study by Zhou et al. (2018)). Figure 1 shows that the concentrations
273 of Cl₂ and ClNO₂ in BASE case are rather low (close to 0), proving that the gas-phase
274 chemistry is not the major pathway to produce Cl₂ and ClNO₂. By contrast, the
275 simulated Cl₂ and ClNO₂ concentrations in HET case increase significantly,
276 correspondingly the NMB and NME changes from -100% to -54% and 100% to 61%
277 for Cl₂, and from -100% to -58% and 100% to 62% for ClNO₂, respectively (the particle
278 surface area concentrations is scaled up by a factor of 5 in daytime and 10 in nighttime
279 because this parameter is underestimated compared to the measured concentrations
280 reported by Zhou et al. (2018)). The simulations of Cl₂ and ClNO₂ are improved as the
281 additional heterogeneous reactions prompt the production of gas phase molecular
282 chlorine. Overall, however, the Cl₂ and ClNO₂ concentrations are still underestimated.
283 Both BASE and HET simulations generally capture the hourly N₂O₅ concentrations as

284 well as the peak values (Figure 1(c)) with similar overall NMB and NME values.

285 The uptake coefficient of O_3 could be an important factor affecting the predicted
286 Cl_2 concentrations as it is found that the heterogeneous reaction of O_3 is the major
287 source of Cl_2 during this period (see discussion in Section 3.2). The influence of
288 different parametrizations of the uptake coefficient of N_2O_5 on $ClNO_2$ and nitrate
289 concentrations are also discussed in Section 3.2.

290 Predicted NO_3^- and PCl concentrations are compared with observations measured
291 at an adjacent monitoring site located at the rooftop of School of Environment building
292 in Tsinghua University (THU, 40.00°N, 116.34°E, about 5 km from IAP) using an
293 Online Analyser of Monitoring of Aerosol and Gases (MARGA) from 11 to 15 June
294 2017. According to Figure 1(d), the simulated nitrate concentration is slightly lower
295 than the observations most of the time. From the evening hours of 12 June to morning
296 hours of 13 June, observed and simulated nitrate concentration both increase
297 significantly. The NMB and NME values of hourly nitrate for the HET case (-5% and
298 39%, respectively) are slightly lower than those for the BASE case -10% and 46%)
299 during this high concentration period. The HET case also generally captures the day-
300 to-day variation of PCl concentration and perform better than the BASE case,
301 correspondingly the NMB and NME are reduced from -48% and 72% to -37% and 67%.
302 The substantial underestimation of PCl in the daytime on 15 June is likely caused by
303 missing local emissions during this period.

304

305 3.2 Impact of uptake coefficients of O_3 and N_2O_5 on chlorine species and nitrate

306 The uptake coefficients of O_3 and N_2O_5 may be important factors affecting the accuracy
307 of simulated nitrate concentrations. Some studies have confirmed that the reaction of
308 O_3 on CPS can indirectly affect the nitrate formation by increasing the atmospheric Cl_2
309 and OH level (Li et al., 2016; Liu et al., 2018). According to Figure 1(a), the improved
310 model still substantially underestimates the concentration of Cl_2 , which may be
311 associated with the underestimation of the uptake coefficients of O_3 , which are
312 empirical and have not been confirmed by laboratory studies. The uptake coefficients
313 were increased by a factor of 10 (0.01 for daytime and 10^{-4} for nighttime) to evaluate

314 the sensitivity of Cl₂ production and nitrate formation to this parameter. Figure 2 shows
315 that the simulated Cl₂ and nitrate concentrations in daytime increase significantly
316 (especially for Cl₂) and sometimes can capture the peak value (such as the daytime peak
317 on 14 June). However, although the NMB and NME of Cl₂ and nitrate improve from -
318 18% and 39% to 1% and 28% when the new uptake coefficients are used, the simulated
319 Cl₂ concentrations are still quite different from the observations (such as during the
320 daytime in 11 and 12 June, see Figure 2). A non-constant parameterization of the uptake
321 coefficients of O₃ that considers the influence of PCl concentrations, meteorology
322 conditions, etc., similar to those of OH and N₂O₅, might be needed. Further laboratory
323 studies should be conducted to provide a better estimation of this important parameter.

324 Several parameterizations for the uptake coefficient of N₂O₅ have been developed
325 for regional and global models and have been evaluated in several previous studies
326 (Tham et al., 2018, McDuffie et al., 2018a, 2018b). In addition to the parameterization
327 of Bertram and Thornton (2009) used in the HET case, two additional simulations were
328 performed to assess the impact of the uptake coefficient of N₂O₅ on nitrate formation.
329 The first simulation uses the original CMAQ parameterization of Davis et al. (2008) and
330 second simulation uses a constant value of 0.09, which is the upper limit of the N₂O₅
331 uptake coefficient derived by Zhou et al. (2018) based on observations. The results from
332 the simulations with the parameterization of Bertram and Thornton (2009) generally
333 agree with the results using those based on Davis et al. (2008). The application of larger
334 and fixed N₂O₅ uptake coefficient leads to slightly better results, which might reflect
335 the fact that the N₂O₅ concentrations are underestimated. Using the uptake coefficient
336 of 0.09 can generally increase the concentration of nitrate in some periods, but it also
337 leads to significant increase of the nitrate level (such as nighttime on 12-13 June and
338 13-14 June), which is 4-6 times higher than those based on Bertram and Thornton
339 (2009). Overall, predicted nitrate concentrations are sensitive to changes in the changes
340 in $\gamma_{N_2O_5}$, with approximately 50% increase in the nitrate when a constant of $\gamma_{N_2O_5}$ of
341 0.09 is used.

342

343 3.3 Spatial distributions of nitrate and chlorine species concentrations

344 The regional distributions of averaged Cl₂, ClNO₂, N₂O₅ and NO₃⁻ concentration from
345 11 to 15 June for the HET case are shown in Figure 3. Compared to the BASE case, the

346 episode average concentrations of Cl_2 and ClNO_2 from the HET case increase
347 significantly in the eastern region of Beijing, reaching up to 23 ppt and 71 ppt from
348 near zero (Figure 3a and 3b). High concentrations are not found in the southern region
349 with intensive emissions of chlorine species (Figure S2). The production of ClNO_2
350 requires the presence of chloride, NO_2 , and O_3 . In the areas close to the fresh emissions,
351 O_3 is generally low (Figure S3), and the production of NO_3 (hence N_2O_5 and ClNO_2) is
352 limited. Therefore, the production rate of ClNO_2 is generally low in areas affected by
353 fresh emissions. Since the contribution of direct emissions to Cl_2 is low and it is
354 predominantly produced secondarily in the atmosphere, high levels of Cl_2 are also
355 found away from the fresh emissions.

356 The spatial distribution of N_2O_5 concentrations differs from that of other species
357 (Figure 3c). While the concentrations of most of the species are higher in the southern
358 region, the N_2O_5 concentrations are lower in some parts of this region. This is because
359 the O_3 concentration in the core urban areas is low due to high NO_x emissions. The
360 N_2O_5 concentrations from the HET case are approximately 16% lower on average
361 (Figure 3d) because the Bertram and Thornton (2009) parameterization used in the HET
362 case generally gives higher uptake coefficients than the parameterization of Davis et al.
363 (2008) used in the BASE case (Table 3).

364 Although the higher uptake coefficients of N_2O_5 in the HET case facilitate faster
365 conversion of N_2O_5 to nitrate, the nitrate concentrations do not always increase. During
366 daytime hours, nitrate concentrations in the HET case increase due to higher OH (Figure
367 3e and Figure 3f, increased OH see Figure S4). At nighttime, in contrast, the nitrate
368 concentration decreases significantly in some regions by about 22%, mainly due to
369 lower molar yield of nitrate from the N_2O_5 heterogeneous reaction in the HET case
370 (Figure 3g and Figure 3h). Although ClNO_2 produced in the N_2O_5 reaction also
371 produces nitrate through a heterogenous reaction when the particle pH is above 2, which
372 is true for most regions (see Figure S5), the uptake coefficient of ClNO_2 is significantly
373 lower than that of N_2O_5 ($0.01\sim 0.09$ for N_2O_5 and 6×10^{-3} for ClNO_2), leading to an
374 overall decrease of nitrate production. As the ClNO_2 production from the heterogeneous
375 reaction leads to less N_2O_5 conversion to non-relative nitrate, it may change the overall

376 lifetime of NO_x and their transport distances. The magnitude of this change and its
377 implications on ozone and PM_{2.5} in local and downwind areas should be further studied.

378

379 3.4 Relationship between nitrate formation and chlorine chemistry

380 Nitrate productions from the homogeneous and heterogeneous pathways in Beijing are
381 approximated by the difference in predicted nitrate concentrations between the BASE
382 or HET case and a sensitivity case without heterogeneous reactions. Averaging over the
383 five-day period, approximately 58% of the nitrate originates from HNO₃ gas-to-particle
384 partitioning and 42% is from heterogeneous reactions (Figure 4). This conclusion
385 generally agrees with measurements at Peking University (PKU) (52% from the
386 heterogeneous process and 48% from HNO₃ partitioning) on four polluted days
387 (average in September 2016 reported by Wang et al. (2017)). Slightly higher
388 contributions of the homogeneous pathway in this study is expected because of high
389 OH concentrations during the day and lower particle surface areas at night.

390 The nitrate formation from different homogeneous and heterogeneous pathways in
391 the BASE case and HET case are further studied. Contributions of different gas phase
392 pathways are determined using the process analysis tool in CMAQ. Contributions of
393 different heterogeneous pathways are determined using a zero-out method that turns off
394 one heterogeneous pathway at a time in a series of sensitivity simulations. Figure 4
395 shows that the reaction of OH and NO₂ is always the major pathway for the formation
396 of nitrate through homogeneous formation of HNO₃ and gas-to-particle partitioning.
397 However, its nitrate production rate through this homogeneous pathway decreases
398 significantly from daytime to nighttime (from 1.81 μg m⁻³ h⁻¹ to 0.33 μg m⁻³ h⁻¹ on
399 average). The nitrate production from other HNO₃ partitioning pathways in the daytime
400 is negligible. At nighttime, homogeneous reaction of N₂O₅ with water vapor accounts
401 for approximately 5% of the overall homogeneous nitrate formation. For the
402 heterogeneous pathways, daytime production rate is approximately 0.6 μg m⁻³ h⁻¹ with
403 1/3 of the contributions from NO₂ and 2/3 from N₂O₅. Nighttime production on nitrate
404 from the heterogeneous pathways is approximately 3.1 μg m⁻³ h⁻¹, of which 85% is due
405 to N₂O₅ and 15% is due to NO₂.

406 Comparing the BASE case and the HET case shows that, when the chlorine
407 chemistry is included, the gaseous HNO_3 produced by OH reacting with NO_2 increases
408 significantly in the HET case. Correspondingly, the nitrate production rate reaches up
409 to $2.04 \mu\text{g m}^{-3} \text{ h}^{-1}$ in the daytime due to increased atmospheric OH concentrations
410 predicted by the chlorine reactions. Similar conclusions are also obtained by Li et al.
411 (2016) and Liu et al. (2017) based on observations and model simulations. The
412 heterogeneous production of nitrate from the reaction of N_2O_5 uptake decreases by
413 approximately 27% in the HET case due to the production of gas phase ClNO_2 .
414 According to the study by Sarwar et al. (2012; 2014), including the heterogeneous
415 reaction of N_2O_5 with PCl decreased the nocturnal nitrate concentration by 11-21% in
416 the United States, which was slightly less than the current study for Beijing. It is likely
417 because PCl concentrations in the United States are significantly lower than those in
418 Beijing (the monthly PCl concentration is $0.06 \mu\text{g m}^{-3}$ in the United State against $\sim 1 \mu\text{g}$
419 m^{-3} in Beijing) so that PCl is depleted quickly. The contributions of NO_2 uptake to
420 nitrate also decrease by 22% because of the lower rate constant of the reaction of NO_2
421 with PCl. In contrast, the contribution of ClNO_2 reacts with particle surface to nitrate
422 production increases by 6% in the HET case. The overall nitrate concentration in the
423 HET case is about 22% higher than that in the BASE case during this study period.

424

425 **4. Conclusions**

426 In this work, a modified CMAQ model incorporated with heterogeneous reactions for
427 the production of molecular chlorine and other reactive chlorine species is developed
428 and its impact on of the nitrate formation predictions are evaluated. The contributions
429 from different homogenous and heterogeneous pathways to nitrate formation are also
430 quantified. High concentration of Cl_2 and ClNO_2 do not occur in the southern part of
431 the Beijing-Tianjin-Hebei region with intensive emissions of chlorine species as higher
432 concentrations of O_3 and N_2O_5 associated with the heterogeneous formation of these
433 species generally occurred in the downwind areas. CTMs without a complete treatment
434 of the chlorine chemistry can underestimate daytime nitrate formation from the
435 homogeneous pathways, particularly from HNO_3 gas-to-particle partitioning due to

436 underestimation of OH concentrations and overestimate the nighttime nitrate formation
437 from the heterogeneous pathways due to missing chlorine heterogeneous chemistry.

438

439 **Data availability.** The data in this study are available from the authors upon request
440 (shxwang@tsinghua.edu.cn)

441

442 **Author contributions.** XQ, QY, SW and JH designed the study; YS, BL, AS, XY
443 provided observation data; XQ, QY, SW, JZ, QX, DD, LD and JX analyzed data. XQ,
444 QY and SW wrote the paper.

445

446 **Competing interests.** The authors declare that they have no conflict of interest.

447

448 **Acknowledgments.** This work was supported by National Natural Science Foundation
449 of China (21625701), China Postdoctoral Science Foundation (2018M641385),
450 National Research Program for Key Issue in Air Pollution Control (DQGG0301,
451 DQGG0501) and National Key R&D Program of China (2018YFC0213805,
452 2018YFC0214006). The simulations were completed on the “Explorer 100” cluster
453 system of Tsinghua National Laboratory for Information Science and Technology.

454 **References**

- 455 Abbatt, J. P., Waschewsky, G. C., et al.: Heterogeneous interactions of HOBr, HNO₃, O₃, and NO₂
456 with deliquescent NaCl aerosols at room temperature. *J. Phys. Chem. A*, 102, 3719-3725, 1998.
- 457 Bertram, T. H., Thornton, J. A.: Toward a general parameterization of N₂O₅ reactivity on aqueous
458 particles: the competing effects of particle liquid water, nitrate and chloride. *Atmos. Chem.*
459 *Phys.*, 9, 8351-8363, 2009.
- 460 Brown, S.S., Stutz, J. Nighttime radical observations and chemistry. *Chemical Society Reviews*, 41,
461 6405-6447, 2012.
- 462 Cater, W. P. L., Heo, G.: Development of revised SAPRC aromatics mechanisms. Final Report to
463 the California Air Resources Board, Contracts No. 07-730 and 08-326, April 12, 2012
- 464 Chang, X., Wang, S., Zhao, B., et al.: Assessment of inter-city transport of particulate matter in the
465 Beijing–Tianjin–Hebei region, *Atmos. Chem. Phys.*, 18, 4843-4858,
466 <https://doi.org/10.5194/acp-18-4843-2018>, 2018.
- 467 Davis, J. M., Bhawe, P. V., Foley, K. M.: Parameterization of N₂O₅ reaction probabilities on the
468 surface of particles containing ammonium, sulfate, and nitrate. *Atmos. Chem. Phys.* 8, 5295-
469 5311, 2008.
- 470 Deng, S., Zhang, C., Liu, Y., et al.: A Full-Scale Field Study on Chlorine Emission of Pulverized
471 Coal-Fired Power Plants in China. *Research of Environmental Science*. In Chinese, 27, 127-
472 133, 2014.
- 473 Deiber, G., George, C., Le Calve, S., Schweitzer, F., Mirabel, P.: Uptake study of ClONO₂ and
474 BrONO₂ by Halide containing droplets. *Atmos. Chem. Phys.* 4, 1291-1299, 2004.
- 475 Ding, D.; Xing, J.; Wang, S. X.; et al.: Emission reductions dominate the decline in the ambient
476 PM_{2.5} concentration and related mortality during 2013-2017 in China. *Environ. Health*
477 *Perspect.*, under review.
- 478 Faxon, C. B., Bean, J. K., Hildebrandt R.L.: Inland Concentrations of Cl₂ and ClONO₂ in Southeast
479 Texas Suggest Chlorine Chemistry Significantly Contributes to Atmospheric Reactivity.
480 *Atmosphere*, 6, 1487-1506, 2015.
- 481 Fu, X., Wang, S.X., Chang, X., et al.: Modeling analysis of secondary inorganic aerosols over China:
482 pollution characteristics, and meteorological and dust impacts. *Sci. Rep.* 6, 35992, 2016.
- 483 Fu, X., Wang, T., Wang, S. X., et al.: Anthropogenic Emissions of Hydrogen Chloride and Fine
484 Particulate Chloride in China. *Environ. Sci. Technol.* 52, 1644-1654, 2018.
- 485 George, I. J., Abbatt, J. P.: Heterogeneous oxidation of atmospheric aerosol particles by gas-phase
486 radicals. *Nat. Chem.* 2, 713-722, 2010.
- 487 Gebel, M. E., Finlayson-Pitts, B. J.: Uptake and reaction of ClONO₂ on NaCl and synthetic sea salt.
488 *J. Phys. Chem. A*, 105, 5178-5187, 2001.
- 489 Glasow, V.R. Wider role for airborne chlorine. *Nature*. 464, 168-169. 2010.
- 490 Goodman, A.L., Underwood, G.M., Grassian, V.H.: Heterogeneous reaction of NO₂:
491 characterization of gas-phase and adsorbed products from the reaction, 2NO₂(g) + H₂O(a) →
492 HONO(g) + HNO₃(a) on Hydrated Silica Particles. *The Journal of Physical Chemistry A* 103,
493 7217-7223, 1999.
- 494 Hu, J., Chen, J., Ying, Q., Zhang, H.: One-Year Simulation of Ozone and Particulate Matter in China
495 Using WRF/CMAQ Modeling System. *Atmos. Chem. Phys.* 16, 10333-10350, 2016.
- 496 Hu, J., Wang, P., Ying, Q., Zhang, H., Chen, J., Ge, X., Li, X., Jiang, J., Wang, S., Zhang, J., Zhao,
497 Y., Zhang, Y.: Modeling biogenic and anthropogenic secondary organic aerosol in China.

498 Atmos. Chem. Phys. 17, 77-92, 2017.

499 Jobson, B. T., Niki, H., Yokouchi, Y., et al.: Measurements of C₂-C₆ hydrocarbons during the Polar
500 Sunrise 1992 experiment: Evidence for Cl atom and Br atom chemistry, *J. Geophys. Res.*, 99,
501 25355-25368, 1994.

502 Keene, W. C.; Pszenny, A. A. P.; Jacob, D. J.; Duce, R. A.; Galloway, J. N.; Schultz-Tokos, J. J.;
503 Sievering, H.; Boatman, J. F.: The Geochemical Cycling of Reactive Chlorine through the
504 Marine Troposphere. *Global. Biogeochem.*, 4, 407-430, 1990.

505 Keene, W. C., Khalil, M. A. K., Erickson, D. J., et al.: Composite global emissions of reactive
506 chlorine from anthropogenic and natural sources: Reactive Chlorine Emissions Inventory. *J.*
507 *Geophys. Res-Atmos.*, 104, 8429-8440, 1999.

508 Kleeman, M.J., Ying, Q., Kaduwela, A.: Control strategies for the reduction of airborne particulate
509 nitrate in California's San Joaquin Valley. *Atmos. Environ.*, 39, 5325-5341, 2005.

510 Knipping, E. M., Lakin, M. J., Foster, K. L., Jungwirth, P., Tobias, D. J., Gerber, R. B., Dabdub, D.,
511 Finlayson-Pitts, B. J.: Experiments and simulations of ion-enhanced interfacial chemistry on
512 aqueous NaCl aerosols. *Science*, 288, 301-306, 2000.

513 Li, H.Y., Zhang, Q., Zheng, B., et al.: Nitrate-driven urban haze pollution during summertime over
514 the North China Plain. *Atmos. Chem. Phys.*, 18, 5293-5306, 2018.

515 Li, Q.Y., Zhang, L., Wang, T., et al. Impacts of heterogeneous uptake of dinitrogen pentoxide and
516 chlorine activation on ozone and reactive nitrogen partitioning: improvement and application
517 of the WRF-Chem model in southern China. *Atmos. Chem. Phys.*, 16, 14875-14890, 2016.

518 Liu, X. X., Qu, H., Huey, L. G., et al.: High Levels of Daytime Molecular Chlorine and Nitryl
519 Chloride at a Rural Site on the North China Plain. *Environ. Sci. Technol.*, 51, 9588-9595, 2017.

520 Liu, Y.M., Fan, Q., Chen, X.Y.: Modeling the impact of chlorine emissions from coal combustion
521 and prescribed waste incineration on tropospheric ozone formation in China. *Atmos. Chem.*
522 *Phys.*, 18, 2709-2724, 2018.

523 Ma, X.Y., Sha, T., Wang, J.Y., et al.: Investigating impact of emission inventories on PM_{2.5}
524 simulations over North China Plain by WRF-Chem. *Atmos. Environ.*, 195, 125-140. 2018.

525 McDuffie, E.E., Fibiger, D.L., Dubé, W.P., et al.: Heterogeneous N₂O₅ uptake during winter: Aircraft
526 measurements during the 2015 WINTER campaign and critical evaluation of current
527 parameterizations. *J. Geophys. Res.: Atmos.*, 123, 4345-4372, 2018a.

528 McDuffie, E.E., Fibiger, D.L., Dubé, W.P., et al.: ClNO₂ yields from aircraft measurements during
529 the 2015 WINTER campaign and critical evaluation of the current parameterization. *J.*
530 *Geophys. Res. Atmos.*, 123, 12-994, 2018b.

531 Osthoff, H.D., Roberts, J.M., Ravishankara, A.R.: High levels of nitryl chloride in the polluted
532 subtropical marine boundary layer. *Nat. Geosci.*, 1, 324-328, 2008.

533 Pratte, P., Rossi, M. J.: The heterogeneous kinetics of HOBr and HOCl on acidified sea salt and
534 model aerosol at 40-90% relative humidity and ambient temperature. *Phys. Chem. Chem. Phys.*
535 8, 3988-4001, 2006.

536 Qiu, X.H., Chai, F.H., Duan, Lei., et al.: Deriving High-Resolution Emission Inventory of Open
537 Biomass Burning in China based on Satellite Observations. *Environ. Sci. Technol.*, 50, 11779-
538 11786, 2016.

539 Riedel, T. P.; Bertram, T. H.; Crisp, T. A.; Williams, E. J.; Lerner, B. M.; Vlasenko, A.; Li, S. M.;
540 Gilman, J.; de Gouw, J.; Bon, D. M.; Wagner, N. L.; Brown, S. S.; Thornton, J. A., Nitryl
541 Chloride and Molecular Chlorine in the Coastal Marine Boundary Layer. *Environ Sci Technol.*,

542 46, 10463-10470, 2012.

543 Roberts, J. M., Osthoff, H. D., Brown, S. S., et al.: N₂O₅ oxidizes chloride to Cl₂ in acidic
544 atmospheric aerosol. *Science*, 321, 1059-1059, 2008.

545 Song, S.J., Gao, M., Xu, W.Q., et al.: Fine-particle pH for Beijing winter haze as inferred from
546 different thermodynamic equilibrium models. *Atmos. Chem. Phys.*, 18, 7423-7438, 2018.

547 Rossi, M J. Heterogeneous Reactions on Salts. *Chemical Reviews*, 103, 4823-4882, 2003.

548 Rudich, Y., Talukdar, R.K., Ravishankara, A.R., et al.: Reactive uptake of NO₃ on pure water and
549 ionic solutions. *J. Geophys. Res.* 101, 21023-21031, 1996.

550 Sarwar, G., Simon, H., Bhawe, P.: Examining the impact of heterogeneous nitryl chloride production
551 on air quality across the United States. *Atmospheric Chemistry and Physics*, 12, 6455-6473,
552 2012.

553 Sarwar, G., Simon, H., Xing, J: Importance of tropospheric ClNO₂ chemistry across the Northern
554 Hemisphere. *Geophysical Research Letters*, 41, 4050-4058, 2014.

555 Seinfeld, J.H., Pandis, S.N.: *Atmospheric Chemistry and Physics: From Air Pollution to Climate*
556 *Change*. Wiley-Interscience, New York.2006.

557 Song, S. J., Gao, M., Xu, W. Q.: Fine-particle pH for Beijing winter haze as inferred from different
558 thermodynamic equilibrium models. *Atmos. Chem. Phys.*, 18, 7423-7438, 2018.

559 Spicer, C.W., Chapman, E.G., Finlayson-Pitts., et al.: Unexpectedly high concentrations of
560 molecular chlorine in coastal air. *Nature*, 394, 353-356, 1998.

561 Thornton, J.A., Kercher, J.P., Riedel, T.P., et al.: A large atomic chlorine source inferred from mid-
562 continental reactive nitrogen chemistry. *Nature*, 464, 271-274, 2010.

563 Tham, Y.J., Wang, Z., Li, Q., et al.: Heterogeneous N₂O₅ uptake coefficient and production yield of
564 ClNO₂ in polluted northern China: roles of aerosol water content and chemical composition.
565 *Atmospheric Chemistry and Physics*, 18, 13155-13171, 2018.

566 Tuazon, E.C., Atkinson, R., Plum, C.N., Winer, A.M., Pitts Jr., J.N.: The reaction of gas phase N₂O₅
567 with water vapor. *Geophysical Research Letters* 10, 953-956, 1983.

568 Wang, S.X., Zhao, B., Cai, S.Y., et al.: Emission trends and mitigation options for air pollutants in
569 East Asia. *Atmos. Chem. Phys.*, 14, 6571-6603, 2014.

570 Wang, H.C., Lu, K.D., Chen, X.R., et al.: High N₂O₅ concentrations observed in urban Beijing:
571 implications of a large nitrate formation pathway. *Environ. Sci. Technol. Lett.* 4, 416-420, 2017.

572 Wen, L., Xue, L.K., Wang, X.F., et al.: Summertime fine particulate nitrate pollution in the North
573 China Plain: increasing trends, formation mechanisms and implications for control policy.
574 *Atmos. Chem. Phys.*, 18, 11261-11275, 2018.

575 Wu, X.W., Chen, W.W., Wang, K., et al.: PM_{2.5} and VOCs emission inventories from cooking in
576 Changchun city. *China Environmental Science*, 38, 2882-2889, 2018.

577 Ying, Q., Li, J. Y., Kota, S. H.: Significant Contributions of Isoprene to Summertime Secondary
578 Organic Aerosol in Eastern United States. *Environ. Sci. Technol.*, 49, 7834-7842, 2015.

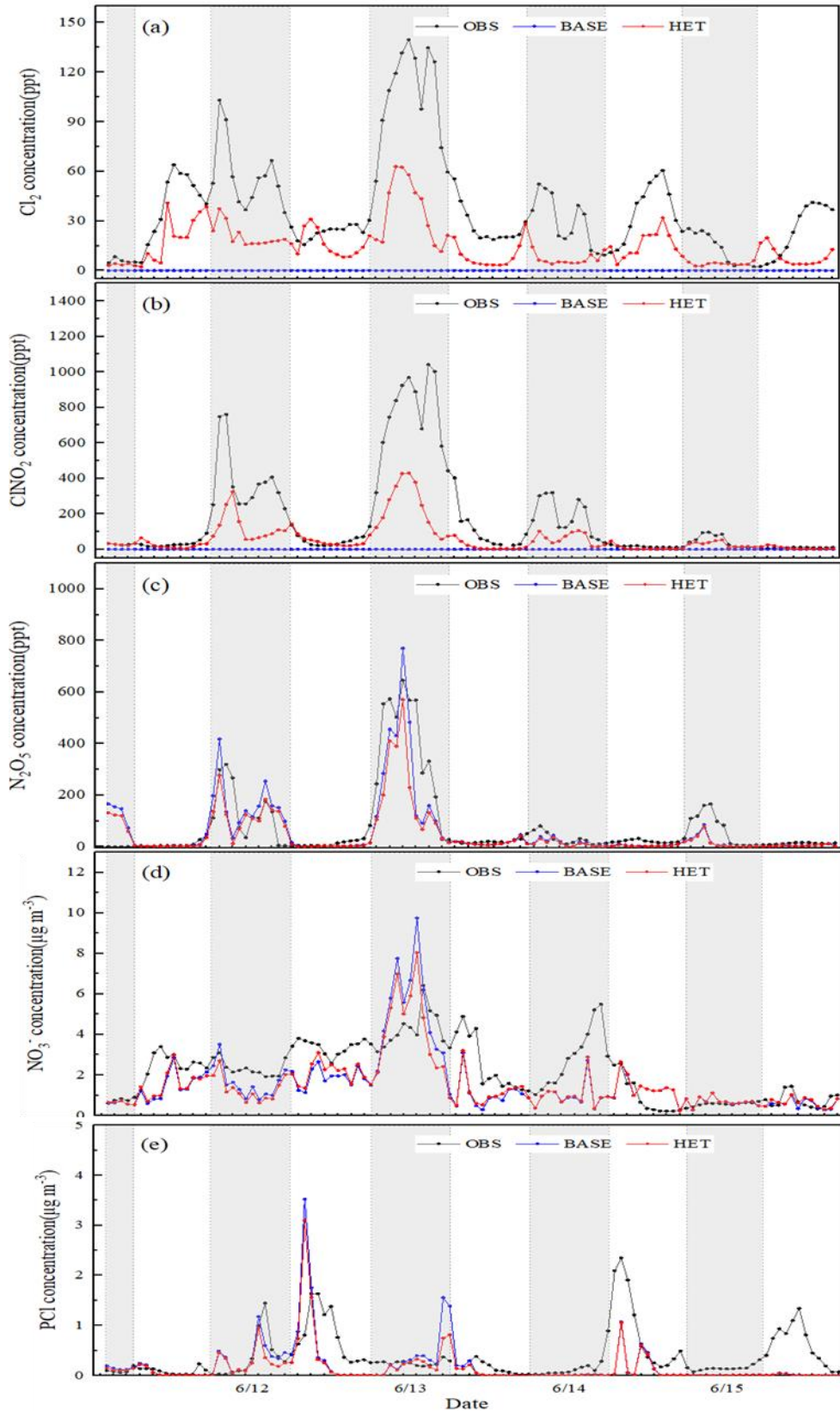
579 Young, C.J., Washenfelder, R.A., Edwards, P.M., et al.: Chlorine as a primary radical: evaluation of
580 methods to understand its role in the initiation of oxidative cycles. *Atmos. Chem. Phys.* 14,
581 3247-3440, 2014.

582 Zhang, T., Peng, L., Li, Y.H.: Chemical characteristics of PM_{2.5} emitted from cooking fumes. *Res.*
583 *Environ. Sci.*, 29, 183-191, 2016. In Chinese.

584 Zhou, W., Zhao, J., Ouyang, B., et al.: Production of N₂O₅ and ClNO₂ in summer in urban Beijing,
585 China., *Atmos. Chem. Phys.*, 18,11581-11597, 2018.

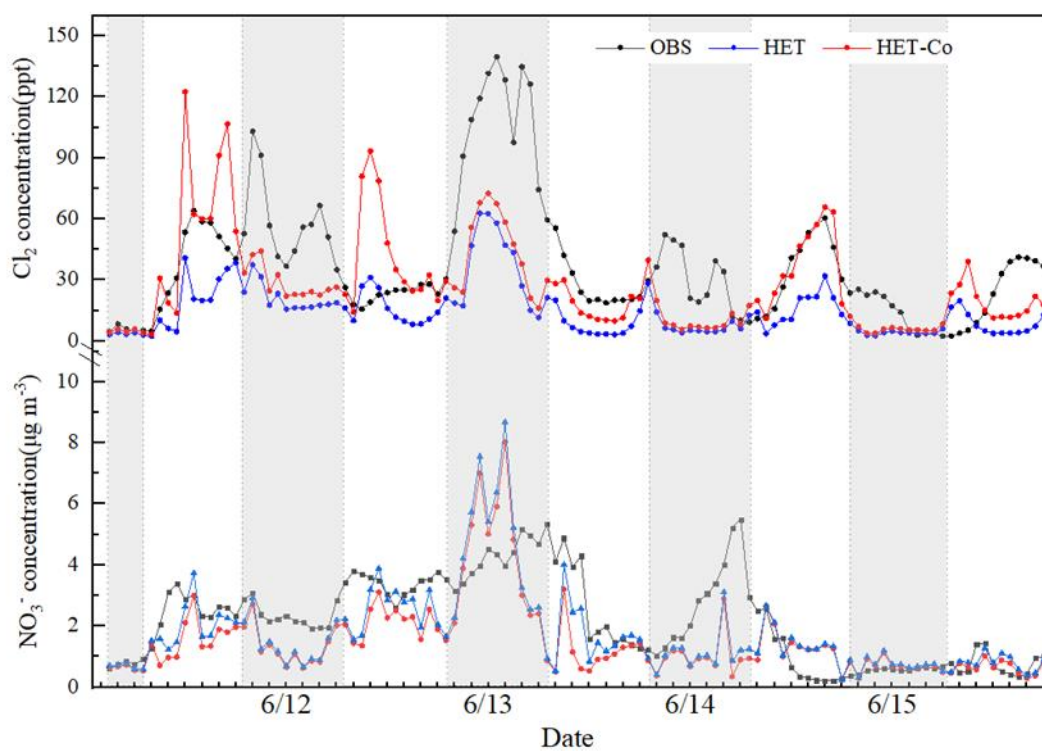
586

587



588

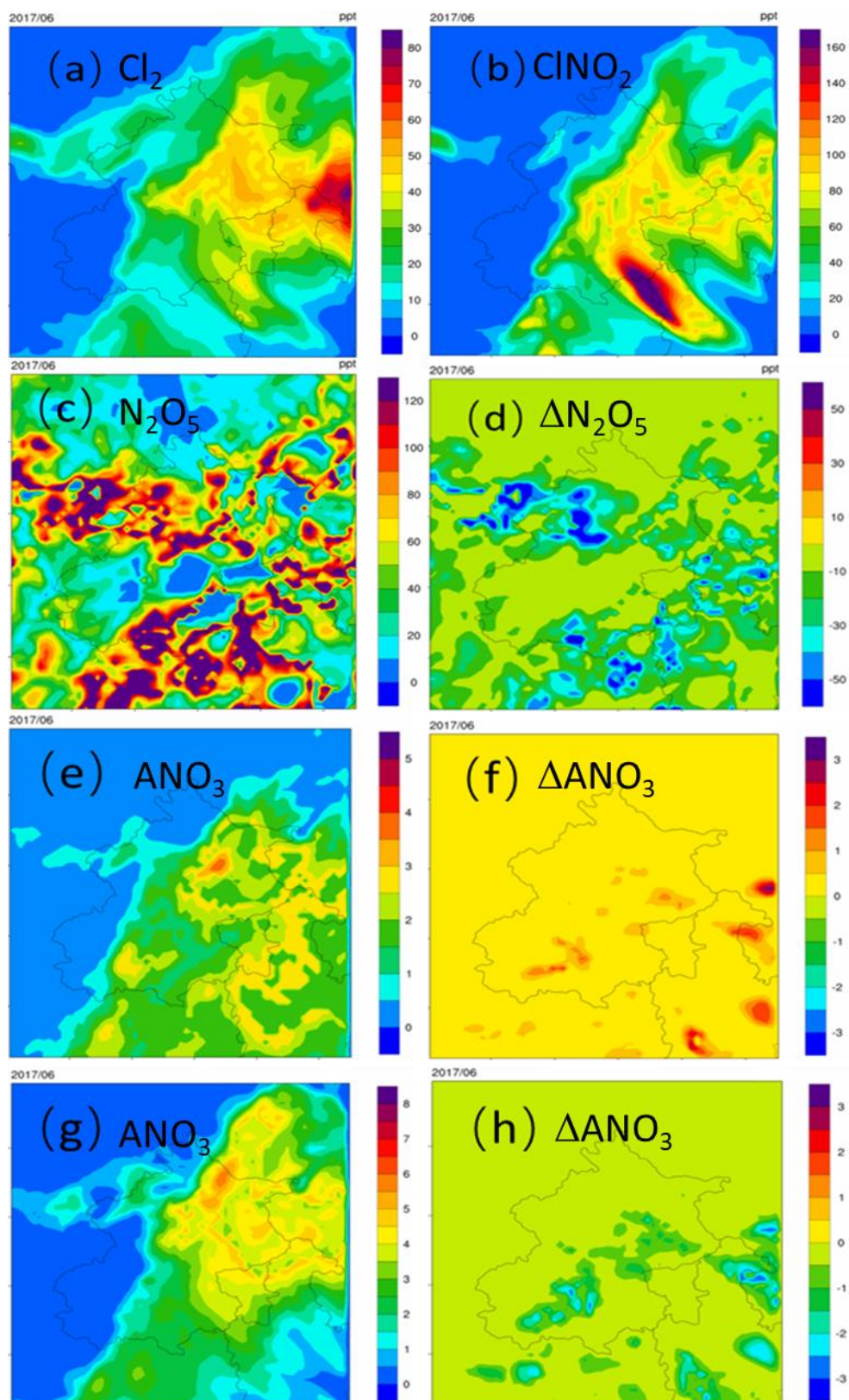
589 Figure 1 Comparison of observed hourly Cl_2 , ClNO_2 , N_2O_5 (at the Institute of
 590 Atmospheric Physics, Chinese Academy of Sciences), NO_3^- and PCl (at Tsinghua
 591 University) in urban Beijing with predictions from the BASE and the HET cases
 592 during 11-15 June 2017.



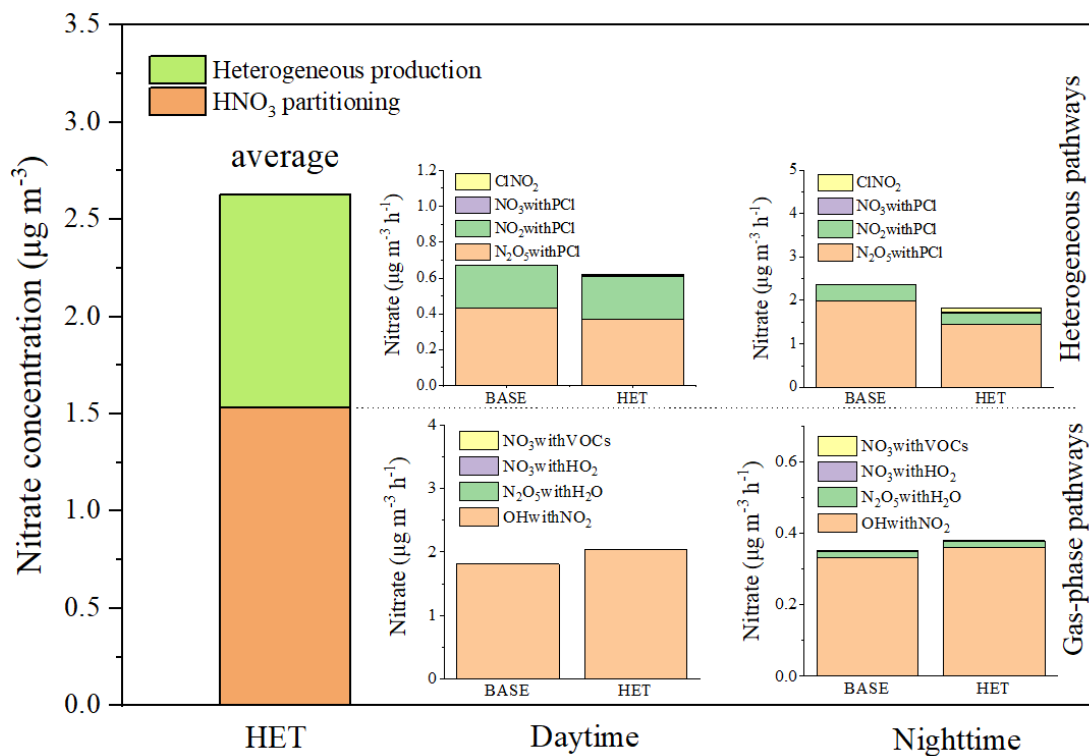
593

594 Figure 2 Comparison of observed and predicted Cl_2 and NO_3^- concentrations under
 595 different uptake coefficient of O_3 (HET: daytime $\gamma_{\text{O}_3} = 1 \times 10^{-3}$, nighttime $\gamma_{\text{O}_3} =$
 596 1×10^{-5} ; HET-Co: daytime $\gamma_{\text{O}_3} = 1 \times 10^{-2}$, nighttime $\gamma_{\text{O}_3} = 1 \times 10^{-4}$).

597



598
 599 Figure 3 Spatial distributions of episode-average (a) Cl_2 , (b) ClNO_2 , (c) N_2O_5 , (e)
 600 daytime nitrate (ANO_3) and (g) nighttime nitrate concentrations from 11-15 June 2017,
 601 and the differences in the episode-average (d) N_2O_5 (HET case – BASE case), (f)
 602 daytime nitrate and (g) nighttime nitrate. Units are $\mu\text{g m}^{-3}$.
 603



604

605 Figure 4 Contributions of different homogeneous and heterogeneous pathways to nitrate
 606 formation.

607

608 Table 1 The sectoral emissions of HCl, Cl₂ and PCl in Beijing in 2017. Unit: Mg year⁻¹
 609 ¹

Sector	Emissions		
	HCl	Cl ₂	PCl
Power plant	22.8	1.2	6.75
Industry	587.3	20.1	89.2
Residential	202.4	8.1	34.7
Biomass burning	0.182	0	0.14
Municipal solid waste	1080.2	0	8.47
Cooking	0	0	426.8
Total	1892.9	29.4	566.1

610

611

612 Table 2 Major gas-phase and heterogeneous pathway of producing nitrate in original CMAQ and newly added or revised heterogeneous reactions
 613 in improved CMAQ.

Type	Reactions	No.	Reference	Comment
Original CMAQ				
Gas-phase chemistry	$\text{OH} + \text{NO}_2 \rightarrow \text{HNO}_3$	R1		
	$\text{N}_2\text{O}_5 + \text{H}_2\text{O} \rightarrow 2\text{HNO}_3$	R7		
	$\text{HO}_2 + \text{NO}_3 \rightarrow 0.2\text{HNO}_3 + 0.8\text{OH} + 0.8\text{NO}_2$	R8		
	$\text{NO}_3 + \text{VOCs}^a \rightarrow \text{HNO}_3$	R9		
Heterogeneous chemistry	$\text{N}_2\text{O}_5(\text{g}) + \text{H}_2\text{O}(\text{aq}) \rightarrow 2\text{H}^+ + 2\text{NO}_3^-$	R5		
	$2\text{NO}_2(\text{g}) + \text{H}_2\text{O}(\text{aq}) \rightarrow \text{HONO}(\text{g}) + \text{H}^+ + \text{NO}_3^-$	R10		
Improved CMAQ				
Newly added or revised heterogeneous reactions	$\text{N}_2\text{O}_5(\text{g}) + \text{H}_2\text{O}(\text{aq}) + \text{Cl}^-(\text{aq}) \rightarrow \text{ClNO}_2(\text{g}) + \text{NO}_3^-$	R6	Bertram and Thornton (2009)	Revise R5
	$2\text{NO}_2(\text{g}) + \text{Cl}^- \rightarrow \text{ClNO}(\text{g}) + \text{NO}_3^-$	R11	Abbatt et al. (1998)	Revise R10
	$\text{NO}_3(\text{g}) + 2\text{Cl}^- \rightarrow \text{Cl}_2(\text{g}) + \text{NO}_3^-$	R12	Rudich et al. (1996)	Increase NO_3^-
	$\text{O}_3(\text{g}) + 2\text{Cl}^- + \text{H}_2\text{O}(\text{aq}) \rightarrow \text{Cl}_2(\text{g}) + \text{O}_2(\text{g}) + 2\text{OH}^-$	R13	Abbatt et al. (1998)	Affect OH
	$2\text{OH}(\text{g}) + 2\text{Cl}^- \rightarrow \text{Cl}_2(\text{g}) + 2\text{OH}^-$	R14	George et al. (2010)	Affect OH
	$\text{ClONO}_2(\text{g}) + \text{Cl}^- \rightarrow \text{Cl}_2(\text{g}) + \text{NO}_3^-$	R15	Deiber et al. (2004)	Affect OH
	$\text{HOCl}(\text{g}) + \text{Cl}^- + \text{H}^+ \rightarrow \text{Cl}_2(\text{g}) + \text{H}_2\text{O}$	R16	Pratte et al. (2006)	Affect OH
	$\text{ClNO}_2(\text{g}) + \text{Cl}^- + \text{H}^+ \rightarrow \text{Cl}_2(\text{g}) + \text{HONO}(\text{g})$ (pH < 2.0)	R17	Riedel et al. (2012)	Affect OH
	$\text{ClNO}_2(\text{g}) + \text{H}_2\text{O}(\text{aq}) \rightarrow \text{Cl}^- + \text{NO}_3^- + 2\text{H}^+$ (pH ≥ 2.0)	R18	Rossi (2003)	Increase NO_3^-

614 ^a: presents different VOCs species. In the SAPRC-11 mechanism, the VOCs species include CCHO (Acetaldehyde), RCHO (Lumped C3+
 615 Aldehydes), GLY (Glyoxal), MGLY (Methyl Glyoxal), PHEN (phenols), BALD (Aromatic aldehydes), MACR (Methacrolein), IPRD (Lumped
 616 isoprene product species).
 617

Table 3 Observed day (D) and night (N) NO_3^- concentrations (Obs.) and predicted uptake coefficient of N_2O_5 ($\gamma_{\text{N}_2\text{O}_5}$) and nitrate concentrations (Pred.) using the parameterizations of $\gamma_{\text{N}_2\text{O}_5}$ by Bertram and Thornton (2009) (Scenario 1), Davis et al., (2008) (Scenario 2) and the upper-limit value derived by Zhou et al. (2018) (Scenario 3)

	NO_3^- Obs.	Scenario1		Scenario2		Scenario3	
		$\gamma_{\text{N}_2\text{O}_5}$	Pred.	$\gamma_{\text{N}_2\text{O}_5}$	Pred.	$\gamma_{\text{N}_2\text{O}_5}$	Pred.
06/11-D	2.54	0.033	1.59	0.008	1.32	0.09	2.17
06/11-12-N	2.42	0.043	1.67	0.037	1.37	0.09	2.12
06/12-D	3.39	0.028	2.16	0.032	2.74	0.09	3.13
06/12-13-N	4.24	0.021	4.02	0.022	4.05	0.09	6.04
06/13-D	2.57	0.012	1.18	0.008	1.06	0.09	2.47
06/13-14-N	4.10	0.022	4.45	0.022	4.45	0.09	7.13
06/14-D	0.95	0.001	1.34	0.001	1.33	0.09	1.64
06/14-15-N	2.75	0.013	1.00	0.007	0.96	0.09	2.33
06/15-D	0.75	0.001	0.66	0.001	0.66	0.09	1.11