

1 **Modeling the impact of heterogeneous reactions of chlorine on**
2 **summertime nitrate formation in Beijing, China**

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19
20 Abstract:

21 A comprehensive chlorine heterogeneous chemistry is incorporated into the
22 Community Multiscale Air Quality (CMAQ) model to evaluate the impact of chlorine-
23 related heterogeneous reaction on diurnal and nocturnal nitrate formation and quantify
24 the nitrate formation from gas-to-particle partitioning of HNO₃ and from different
25 heterogeneous pathways. The results show that these heterogeneous reactions increase
26 the atmospheric Cl₂ and ClNO₂ level (~100%), which further affect the nitrate
27 formation. Sensitivity analyses of uptake coefficients show that the empirical uptake
28 coefficient for the O₃ heterogeneous reaction with chlorinated particles may lead to the
29 large uncertainties in the predicted Cl₂ and nitrate concentrations. The N₂O₅ uptake

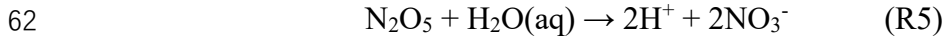
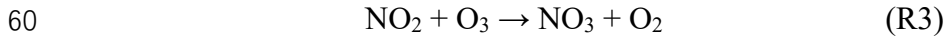
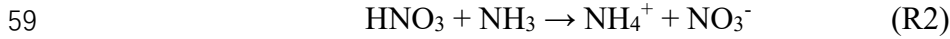
30 coefficient with particulate Cl⁻ concentration dependence performs better to capture the
31 concentration of ClNO₂ and nocturnal nitrate concentration. The reaction of OH and
32 NO₂ in daytime increases the nitrate by ~15% when the heterogeneous chlorine
33 chemistry is incorporated, resulting in more nitrate formation from HNO₃ gas-to-
34 particle partitioning. By contrast, the contribution of the heterogeneous reaction of
35 N₂O₅ to nitrate concentrations decreases by about 27% in the nighttime when its
36 reactions with chlorinated particles are considered. However, the generated gas-phase
37 ClNO₂ from the heterogeneous reaction of N₂O₅ and chlorine-containing particles
38 further reacts with the particle surface to increase the nitrate by 6%. In general, this
39 study highlights the potential of significant underestimation of daytime and
40 overestimation of nighttime nitrate concentrations for chemical transport models
41 without proper chlorine chemistry in the gas and particle phases.

42

43 **1. Introduction**

44 In recent years, nitrate has become the primary component of PM_{2.5} (particulate matter
45 with aerodynamic diameter less than 2.5μm) in Beijing with sustained and rapid
46 reduction of SO₂ and primary particulate matter emissions (Ma et al., 2018; Li et al.,
47 2018; Wen et al., 2018). Observations showed that the relative contributions of
48 secondary nitrate in PM_{2.5} could reach up to approximately 50% during some severe
49 haze pollution days (Li et al., 2018). The mechanism of secondary nitrate formation can
50 be summarized as two major pathways: (1) Gas-to-particle partitioning of HNO₃, which
51 happens mostly in the daytime. The reaction of OH with NO₂ produces gaseous HNO₃,
52 which subsequently partition into the particle phase. The existence of NH₃ or basic
53 particles enhances this process by NH₃-NH₄⁺ gas-particle equilibrium (Kleeman et al.,
54 2005; Seinfeld and Pandis, 2006); (2) Hydrolysis of N₂O₅, which is more important at
55 nighttime. N₂O₅ forms from the reactions of NO₂, O₃ and NO₃ and hydrolyzes to
56 produce particulate nitrate. They can be summarized as reactions R1-R5 (Brown and
57 Stutz 2012):

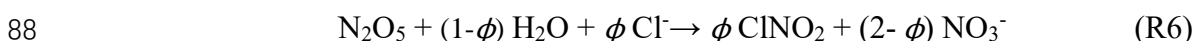




63 In addition to reactions R1 and R5, gas phase reactions of NO_3 with HO_2 and VOCs,
64 N_2O_5 with water vapor (Tuazon et al., 1983) and the heterogeneous reaction of NO_2
65 with water-containing particle (Goodman et al., 1999) produce HNO_3 or nitrate as well.
66 These reactions are listed in Table 2 as reactions R8, R9 and R10.

67 However, chemistry transport models (CTMs) incorporated with these mechanisms
68 still can't accurately capture the spatiotemporal distributions of nitrate in some studies
69 in polluted northern China. For example, Chang et al. (2018) showed that the simulated
70 nitrate concentrations derived from the default CMAQ (version 5.0.2) were
71 significantly higher than the observations in summer at two sites adjacent to Beijing.
72 Fu et al. (2016) also found that default CMAQ (version 5.0.1) overestimated the
73 simulated nitrate concentrations in the Beijing-Tianjin-Hebei region.

74 In recent fields studies, it was found that high particulate chlorine emissions might
75 have a significant impact on the oxidation capacity of the urban atmosphere and thus
76 could affect nitrate concentrations. According to the field measurements in June 2017
77 in Beijing (Zhou et al., 2018), the 2-min averaged concentrations of reactive molecular
78 chlorine (Cl_2) and nitryl chloride (ClNO_2) reached up to 1000 pptv and 1200 pptv,
79 respectively, during some severe air pollution periods in summer. The Cl_2
80 concentrations were significantly higher than those observed in North American coastal
81 cities affected by onshore flow and the lower atmosphere in the remote Arctic region
82 (Spicer et al., 1998; Glasow et al., 2010; Liu et al., 2017). During these pollution events,
83 the corresponding concentrations of N_2O_5 (2-min average) and nitrate (5-min average)
84 rose from 40 pptv and $1 \mu\text{g m}^{-3}$ to 700 pptv and $5 \mu\text{g m}^{-3}$. To explain the high levels of
85 ClNO_2 , some studies suggested that reaction R5 should be revised to account for ClNO_2
86 production from the heterogeneous reaction of N_2O_5 on chloride-containing particles
87 (CPS) (Osthoff et al., 2008; Thornton et al., 2010), as shown in reaction R6:



89 where ϕ represents the molar yield of ClNO₂. When Cl⁻ is enough, this reaction leads
90 to lower nitrate concentrations than reaction R5. By incorporating this reaction into
91 WRF-Chem, Li et al. (2016) found that the improved model performed better to match
92 the observed nitrate concentrations in Hongkong during 15 November and 5 December
93 2013. However, ClNO₂ could affect the formation of nitrate indirectly by increasing the
94 atmospheric OH after a series of chemical reactions, which are briefly summarized into
95 three steps: (1) the photolysis of ClNO₂ produces chlorine radicals (Cl[•]); (2) the reaction
96 of Cl[•] with VOCs produces peroxy radical (HO₂ and RO₂); and (3) the increased HO₂
97 and RO₂ prompt the formation of OH through HO_x cycle and lead to increased HNO₃
98 production (Young et al., 2014; Jobson et al., 1994). The overall impact of R6 on nitrate
99 remains to be investigated.

100 Another related but unresolved issue is the sources of the high concentrations Cl₂,
101 which could not be explained by the N₂O₅ heterogeneous reaction with Cl⁻ and the
102 subsequent reactions of ClNO₂ in the gas phase. It has been reported that the reactions
103 of gaseous O₃, OH, HO₂, ClNO₂, hypochlorous acid (HOCl), chlorine nitrate (ClONO₂)
104 with CPS can produce Cl₂, which can subsequently photolyze to produce Cl[•]
105 (Knipping et al., 2000; George et al., 2010; Pratte et al., 2006; Deiber et al., 2004; Faxon
106 et al., 2015). However, these heterogeneous reactions on CPS are generally missing in
107 most of the current CTMs and it is unclear whether these reactions will be able to
108 explain the observed Cl₂ concentrations and the overall impact of these reactions on
109 nitrate.

110 Previously, biomass burning, coal combustion, and waste incineration were
111 identified as the main sources of gaseous and particulate chlorine compounds in China
112 from International Global Atmospheric Chemistry Program's Global Emissions
113 Inventory Activity (GEIA) based on the year 1990 and a localized study by Fu et al.
114 based on the year 2014 (Keene et al., 1999; Fu et al., 2018). However, recent source
115 apportionment results of PM_{2.5} in Beijing showed that the contribution of coal
116 combustion had extremely decreased from 22.4% in 2014 to 3% in 2017 with the
117 replacement of natural gas (obtained from official website of Beijing Municipal Bureau

118 of Statistics, available at <http://edu.bjstats.gov.cn/>). Another important source—cooking
119 has received attention as its increasing contribution to PM_{2.5} (accounting for 33% of the
120 residential sector; obtained from the official source apportionment analysis of PM_{2.5} in
121 Beijing in 2017; see <http://www.bjepb.gov.cn/bjhr-b/index/index.html>). Moreover, the
122 high content of particulate sodium chloride was measured from the source
123 characterization studies of PM_{2.5} released from the cooking activities (Zhang et al.,
124 2016). Thus, it is necessary to compile an updated emission inventory for Beijing to
125 include the emissions from cooking and other sources (coal burning, solid waste
126 incineration, biomass burning, etc.) in order to explore the emissions of the chlorine
127 species on atmospheric nitrate formation.

128 In this study, a CMAQ model with an improved chlorine heterogeneous chemistry
129 is applied to simulate summer nitrate concentration in Beijing. Sensitivity simulations
130 are conducted to evaluate the contributions of HNO₃ gas-to-particle partitioning and
131 heterogenous production to aerosol nitrate. The results of this work can improve our
132 understandings on nitrate formation and provide useful information on nitrate pollution
133 control strategies in Beijing.

134

135 **2. Emissions, chemical reactions and model description**

136 2.1 Emissions

137 Generally, the conventional emission inventories of air pollutants in China only include
138 the common chemical species, such as SO₂, NO_x, VOCs, PM_{2.5}, PM₁₀, NH₃, BC, and
139 OC (Wang et al., 2014). Chloride compound emissions were not included. However,
140 the emissions of chlorine species are vital for studying the chlorine chemical
141 mechanism. Recently, the inorganic hydrogen chloride (HCl) and fine particulate
142 chloride (PCI) emission inventories for the sectors of coal combustion, biomass burning,
143 and waste incineration were developed for the year of 2014 (Qiu et al., 2016, Fu et al.,
144 2018, Liu et al., 2018). However, the gaseous chlorine emission was not estimated in
145 these studies. In addition, these studies did not account for the rapid decrease of coal
146 consumption in recent years in Beijing, from 2000 Mt in 2014 to 490 Mt in 2017. More

147 importantly, the cooking source, as one of the major contributors to particulate chlorine
148 in Beijing, is not included in current chlorine emission inventories. Thus, a new
149 emission inventory of reactive chlorine species, which includes HCl, Cl₂ and PCl, were
150 developed in this study for the year of 2017.

151 The emission factor method (equation (1)) is applied to calculate the emissions of
152 these reactive chlorine species from coal combustion, biomass burning, municipal solid
153 waste incineration and industrial processes:

$$154 \quad E_{i,j} = A_i \times EF_{i,j} \quad (1)$$

155 where $E_{i,j}$ represents the emission factor of pollutant j in sector i ; A represents the
156 activity data; EF represents the emission factor. EF for PCl is estimated by $EF_{i,PCl} =$
157 $EF_{i,PM_{2.5}} \times f_{Cl,i}$, where $f_{Cl,i}$ represents the mass fraction of PCl in primary PM_{2.5}.
158 Activity data are obtained from the Beijing Municipal Bureau of Statistics (available at
159 <http://tjj.beijing.gov.cn/>). The Cl₂ emission factor for coal combustion is calculated
160 based on the content of Cl in coal, which had been measured by Deng et al (2017). The
161 PM_{2.5} emission factors and mass fractions of PCl in PM_{2.5} to calculate the emissions of
162 Cl had been described in detail by Fu et al. (2018). PCl in PM_{2.5} for coal combustion
163 and biomass burning are taken as 1% and 9.0%, respectively, based on local
164 measurements in Beijing.

165 Emissions of PCl from cooking, including contributions from commercial and
166 household cooking, are estimated using equation (2):

$$167 \quad E_{PCl} = [N_f \times V_f \times H_f \times EF_{f,PCl} + V_c \times H_c \times N_c \times n \times EF_{c,PCl} \times (1 - \eta)] \times 365 \quad (2)$$

168 where N_f is the number of households, V_f is the volume of exhaust gas from a household
169 stove (2000 m³ h⁻¹); H_f is the cooking time for a family (0.5 h day⁻¹); $EF_{f,PCl}$ and
170 $EF_{c,PCl}$ are the emission factors (kg m⁻³) of PCl for household and commercial cooking,
171 respectively; H_c is the cooking time in a commercial cooking facility (6 h day⁻¹); N_c is
172 the number of restaurants, schools and government departments. V_c is the volume of
173 exhaust gas from a commercial cooking stove (8000 m³ h⁻¹); n is the number of stoves
174 for each unit, which equals to 6 for a restaurant and is calculated as one stove per 150
175 students for each school. η is the removal efficiency of fume scrubbers (30%). $EF_{c,PCl}$
176 is the emission factor (kg m⁻³) of PCl in commercial cooking. These constants are all

177 based on Wu et al. (2018). The PCl fraction in PM_{2.5} from cooking is take as 10%, based
178 on local measurements. HCl and Cl₂ emissions from cooking are not considered in this
179 study.

180 The sectoral emissions of HCl, Cl₂ and PCl are summarized in Table 1. The
181 estimated HCl, Cl₂ and PCl emissions in Beijing are 1.89 Gg, 0.07Gg and 0.63Gg
182 respectively. The Cl emissions estimated for 2014 by Fu et al. (2018) were used for
183 other areas. This simplification is a good approximation because replacing coal with
184 natural gas only occurred in Beijing, and reduction of coal consumption in surrounding
185 regions was generally less than 15%. In addition, strict control measures for biomass
186 burning, cooking and municipal solid waste incineration have not been implemented in
187 most regions yet. Emissions of conventional species for this study period are developed
188 in a separate study that is currently under review and are summarized in Table S1.

189

190 2.2 Chlorine-related heterogeneous reactions

191 The heterogeneous reactions in original CMAQ (version 5.0.1) are not related to
192 chlorine species. In this study, the original heterogeneous reactions of N₂O₅ and NO₂
193 (R5 and R10 in Table 2) are replaced with a revised version which includes production
194 of ClNO₂ from CPS (R6 and R11 in Table 2). In reaction R6, the molar yield of ClNO₂
195 (ϕ_{ClNO_2}) is represented as equation (3) (Bertram and Thornton, 2009):

$$\phi_{ClNO_2} = \left(1 + \frac{[H_2O]}{483 \times [Cl^-]} \right)^{-1} \quad (3)$$

196 where [H₂O] and [Cl⁻] are the molarities of liquid water and chloride (mol m⁻³).

197 In addition, laboratory observations confirmed that the heterogeneous uptake of
198 some oxidants (such as O₃ and OH) and reactive chlorine species (such as ClNO₂, HOCl,
199 and ClONO₂) could also occur on CPS to produce Cl₂. These reactions are implemented
200 in the model and summarized in Table 2 as R13-R18. Note that the products from the
201 heterogeneous uptake of ClNO₂ on CPS vary with particle acidity (Riedel et al., 2012;
202 Rossi, 2003). It generates Cl₂ under the condition of pH lower than 2 but produces
203 nitrate and chloride under higher pH conditions. The reaction rates of the heterogeneous

204 reactions are parameterized as first-order reactions, with the rate of change of gas phase
 205 species concentrations determined by equations (4) (Ying et al., 2015):

$$\frac{dC}{dt} = -\frac{1}{4}(v\gamma A)C = -k^I C \quad (4)$$

206 where C represents the concentration of species; v represents the thermal velocity of
 207 the gas molecules (m s^{-1}); A is the CMAQ-predicted wet aerosol surface area
 208 concentration ($\text{m}^2 \text{m}^{-3}$); γ represents the uptake coefficient. For all gas phases species
 209 (except ClNO_2) involved in the heterogeneous reactions (R6 and R11-R18), a simple
 210 analytical solution can be used to update their concentrations from time t_0 to $t_0+\Delta t$:
 211 $[C]_{t_0+\Delta t}=[C]_{t_0} \exp(-k^I \Delta t)$, where Δt is the operator-splitting time step for heterogeneous
 212 reactions.

213 The rate of change of ClNO_2 includes both removal and production terms, as
 214 shown in equation (5):

$$\begin{aligned} \frac{d[\text{ClNO}_2]}{dt} &= -k_i^I [\text{ClNO}_2] + k_6^I \phi_{\text{ClNO}_2} [\text{N}_2\text{O}_5] \\ &= -k_i^I [\text{ClNO}_2] + k_6^I \phi_{\text{ClNO}_2} [\text{N}_2\text{O}_5]_{t_0} \exp(-k_6^I t) \end{aligned} \quad (5)$$

215 Assuming ϕ_{ClNO_2} is a constant, an analytical solution for equation (5) can be obtained,
 216 as shown in equation (6):

$$\begin{aligned} [\text{ClNO}_2]_{t_0+\Delta t} &= [\text{ClNO}_2]_{t_0} \exp(-k_i^I \Delta t) \\ &+ \frac{k_6^I \phi_{\text{ClNO}_2} [\text{N}_2\text{O}_5]_{t_0}}{k_i^I - k_6^I} [\exp(-k_6^I \Delta t) - \exp(-k_i^I \Delta t)] \end{aligned} \quad (6)$$

217 where k_i^I represents the pseudo first-order rate coefficient of either reaction R17 or
 218 R18, depending on pH.

219 The uptake coefficients γ of gaseous species are obtained from published
 220 laboratorial studies. In the original CMAQ, the uptake coefficient of N_2O_5 is determined
 221 as a function of the concentrations of $(\text{NH}_4)_2\text{SO}_4$, NH_4HSO_4 and NH_4NO_3 (Davis et al.,
 222 2008). In this study, the $\text{P}(\text{Cl})$ and NO_3^- dependent parameterization by Bertram and
 223 Thornton (2009) (equation (7)) is used:

$$\gamma_{\text{N}_2\text{O}_5} = \begin{cases} 0.02, & \text{for frozen aerosols} \\ 3.2 \times 10^{-8} K_f \left[1 - \left(1 + \frac{6 \times 10^{-2} [\text{H}_2\text{O}]}{[\text{NO}_3^-]} + \frac{29 [\text{Cl}^-]}{[\text{NO}_3^-]} \right)^{-1} \right] & \end{cases} \quad (7)$$

224 In the above equation, K_f is parameterized function based on molarity of water: $K_f =$

225 $1.15 \times 10^6(1 - e^{-0.13[\text{H}_2\text{O}]})$. NO_3^- and Cl^- concentrations are also in molarity. The
226 uptake coefficient of OH is expressed in equation (8) as a function of the concentration
227 of PCl following the IUPAC (International Union of Pure and Applied Chemistry,
228 available at [http://iupac.poleether.fr/htdocs/datasheets/pdf/O-](http://iupac.poleether.fr/htdocs/datasheets/pdf/O-H_halide_solutions_VI.A2.1.pdf)
229 [H_halide_solutions_VI.A2.1.pdf](http://iupac.poleether.fr/htdocs/datasheets/pdf/O-H_halide_solutions_VI.A2.1.pdf)).

$$\gamma = \min\left(0.04 \times \frac{[\text{Cl}^-]}{1000 \times M}, 1\right) \quad (8)$$

230 where M represents the volume of liquid water in aerosol volume ($\text{m}^3 \text{m}^{-3}$). For frozen
231 particles, the uptake coefficient is limited to 0.02, as used in the original CMAQ model.

232 The uptake coefficients of O_3 , NO_3 , NO_2 , HOCl, ClNO₂, and ClONO₂ are treated
233 as constants. Among of them, the γ values of NO_3 , NO_2 , HOCl and ClONO₂ are set as
234 3×10^{-3} , 1×10^{-4} , 1.09×10^{-3} and 0.16 based on laboratory measurements (Rudich et al.,
235 1996; Abbatt et al., 1998; Pratte et al., 2006; Gebel et al., 2001). A preliminary value
236 of 10^{-3} in the daytime and 10^{-5} at nighttime is chosen for the O_3 uptake coefficient. The
237 daytime γ_{O_3} is based on the analysis of Cl_2 production rate in a hypothesized
238 geochemical cycle of reactive inorganic chlorine in the marine boundary layer by Keene
239 et al. (1990). The lower nighttime value was also recommended by Keene et al. (1990)
240 who noted that Cl_2 production in the marine boundary layer are lower at night. The
241 uptake coefficient of ClNO₂ depends on the particle acidity, with the value of 2.65×10^{-6}
242 for reaction R17 and 6×10^{-3} for reaction R18 (Robert et al., 2008).

243

244 2.4 CMAQ model configuration

245 These heterogeneous reactions of chlorine are incorporated into a revised CMAQ based
246 on the CMAQ version 5.0.1 to simulate the distribution of nitrate concentration in
247 Beijing from 11 to 15 June 2017. The revised CMAQ model without heterogeneous
248 reactions of chlorine has been described in detail by Ying et al. (2015) and Hu et al.
249 (2016, 2017). In summary, the gas phase chemical mechanism in the revised CMAQ
250 model is based on the SAPRC-11 (Cater et al., 2012) with a comprehensive inorganic
251 chlorine chemistry. Reactions of Cl radical with several major VOCs, which lead to
252 production of HCl, are also included. The aerosol module is based on AERO6 with an

253 updated treatment of NO₂ and SO₂ heterogeneous reaction and formation of secondary
254 organic aerosol from isoprene epoxides. Three-level nested domains with the
255 resolutions of 36km, 12km, and 4km using Lambert Conformal Conic projection
256 (173×136, 135×228 and 60×66 grid cells) are chosen in this work (the domains see
257 Figure S1). The two true latitudes are set to 25°N and 40°N and the origin of the domain
258 is set at 34°N, 110°E. The left-bottom coordinates of the outmost domain are positioned
259 at x = -3114 km, y = -2448 km. The BASE case (heterogeneous reactions of Cl turned
260 off) and HET case (all heterogeneous reactions enabled) are compared to evaluate the
261 impact of heterogeneous chlorine chemistry on nitrate formation.

262

263 **3. Results**

264 3.1 Model performance evaluation

265 Predicted O₃, NO₂ and PM_{2.5} concentrations from the BASE case simulation are
266 evaluated against monitoring data at 12 sites in Beijing (Table S2) in 11 to 15, June
267 2017. The average NMB/NME values for O₃, NO₂ and PM_{2.5} across the 12 sites are -
268 8%/29%, -7%/59% and -8%/53%, respectively. Predicted hourly Cl₂, ClNO₂ and N₂O₅
269 concentrations were compared with observations measured at the Institute of
270 Atmospheric Physics (IAP), Chinese Academy of Sciences (39.98°N, 116.37°E) using
271 a high-resolution time-of-flight chemical ionization mass spectrometer (CIMS) from 11
272 to 15 June 2017 (for site description, instrument introduction, and analytical method,
273 please refer to the study by Zhou et al. (2018)). Figure 1 shows that the concentrations
274 of Cl₂ and ClNO₂ in BASE case are rather low (close to 0), proving that the gas-phase
275 chemistry is not the major pathway to produce Cl₂ and ClNO₂. By contrast, the
276 simulated Cl₂ and ClNO₂ concentrations in HET case increase significantly,
277 correspondingly the NMB and NME changes from -100% to -54% and 100% to 61%
278 for Cl₂, and from -100% to -58% and 100% to 62% for ClNO₂, respectively (the particle
279 surface area concentrations is scaled up by a factor of 5 in daytime and 10 in nighttime
280 because this parameter is underestimated compared to the measured concentrations
281 reported by Zhou et al. (2018)). The simulations of Cl₂ and ClNO₂ are improved as the

282 additional heterogeneous reactions prompt the production of gas phase molecular
283 chlorine. Overall, however, the Cl_2 and ClNO_2 concentrations are still underestimated.
284 Both BASE and HET simulations generally capture the hourly N_2O_5 concentrations as
285 well as the peak values (Figure 1(c)) with similar overall NMB and NME values.

286 The uptake coefficient of O_3 could be an important factor affecting the predicted
287 Cl_2 concentrations as it is found that the heterogeneous reaction of O_3 is the major
288 source of Cl_2 during this period (see discussion in Section 3.2). The influence of
289 different parametrizations of the uptake coefficient of N_2O_5 on ClNO_2 and nitrate
290 concentrations are also discussed in Section 3.2.

291 Predicted NO_3^- and PcI concentrations are compared with observations measured
292 at an adjacent monitoring site located at the rooftop of School of Environment building
293 in Tsinghua University (THU, 40.00°N, 116.34°E, about 5 km from IAP) using an
294 Online Analyser of Monitoring of Aerosol and Gases (MARGA) from 11 to 15 June
295 2017. According to Figure 1(d), the simulated nitrate concentration is slightly lower
296 than the observations most of the time. From the evening hours of 12 June to morning
297 hours of 13 June, observed and simulated nitrate concentration both increase
298 significantly. The NMB and NME values of hourly nitrate for the HET case (-5% and
299 39%, respectively) are slightly lower than those for the BASE case (-10% and 46%)
300 during this high concentration period. The HET case also generally captures the day-
301 to-day variation of PcI concentration and perform better than the BASE case,
302 correspondingly the NMB and NME are reduced from -48% and 72% to -37% and 67%.
303 The substantial underestimation of PcI in the daytime on 15 June is likely caused by
304 missing local emissions during this period.

305

306 3.2 Impact of uptake coefficients of O_3 and N_2O_5 on chlorine species and nitrate

307 The uptake coefficients of O_3 and N_2O_5 may be important factors affecting the accuracy
308 of simulated nitrate concentrations. Some studies have confirmed that the reaction of
309 O_3 on CPS can indirectly affect the nitrate formation by increasing the atmospheric Cl_2
310 and OH level (Li et al., 2016; Liu et al., 2018). According to Figure 1(a), the improved
311 model still substantially underestimates the concentration of Cl_2 , which may be

312 associated with the underestimation of the uptake coefficients of O₃, which are
313 empirical and have not been confirmed by laboratory studies. The uptake coefficients
314 were increased by a factor of 10 (0.01 for daytime and 10⁻⁴ for nighttime) to evaluate
315 the sensitivity of Cl₂ production and nitrate formation to this parameter. Figure 2 shows
316 that the simulated Cl₂ and nitrate concentrations in daytime increase significantly
317 (especially for Cl₂) and sometimes can capture the peak value (such as the daytime peak
318 on 14 June). However, although the NMB and NME of Cl₂ and nitrate improve from -
319 18% and 39% to 1% and 28% when the new uptake coefficients are used, the simulated
320 Cl₂ concentrations are still quite different from the observations such as during the
321 daytime in 11 and 12 June, see Figure 2). A non-constant parameterization of the uptake
322 coefficients of O₃ that considers the influence of PCl concentrations, meteorology
323 conditions, etc., similar to those of OH and N₂O₅, might be needed. Further laboratory
324 studies should be conducted to provide a better estimation of this important parameter.

325 Several parameterizations for the uptake coefficient of N₂O₅ have been developed
326 for regional and global models and have been evaluated in several previous studies
327 (Tham et al., 2019, McDuffie et al., 2018a, 2018b). In addition to the parameterization
328 of Bertram and Thornton (2009) used in the HET case, two additional simulations were
329 performed to assess the impact of the uptake coefficient of N₂O₅ on nitrate formation.
330 The first simulation uses the original CMAQ parameterization of Davis et al.(2008) and
331 second simulation uses a constant value of 0.09, which is the upper limit of the N₂O₅
332 uptake coefficient derived by Zhou et al. (2018) based on observations. The results from
333 the simulations with the parameterization of Bertram and Thornton (2009) generally
334 agree with the results using those based on Davis et al. (2008) and the larger constant
335 N₂O₅ leads to slightly better results, which might reflect the fact that the N₂O₅
336 concentrations are underestimated. Using the uptake coefficient of 0.09 can generally
337 increase the concentration of nitrate in some periods, but it also leads to significant
338 increase of the nitrate level (such as nighttime on 12-13 June and 13-14 June), which is
339 4-6 times higher than those based on Bertram and Thornton (2009). Overall, predicted
340 nitrate concentrations are sensitive to changes in the changes in $\gamma_{N_2O_5}$, with
341 approximately 50% increase in the nitrate when a constant of $\gamma_{N_2O_5}$ of 0.09 is used.

342

343 3.3 Spatial distributions of nitrate and chlorine species concentrations

344 The regional distributions of averaged Cl_2 , ClNO_2 , N_2O_5 and NO_3^- concentration from
345 11 to 15 June for the HET case are shown in Figure 3. Compared to the BASE case, the
346 episode average concentrations of Cl_2 and ClNO_2 from the HET case increase
347 significantly in the eastern region of Beijing, reaching up to 23 ppt and 71 ppt from
348 near zero (Figure 3a and 3b). High concentrations are not found in the southern region
349 with intensive emissions of chlorine species (Figure S2). The production of ClNO_2
350 requires the presence of chloride, NO_2 , and O_3 . In the areas close to the fresh emissions,
351 O_3 is generally low (Figure S3), and the production of NO_3 (hence N_2O_5 and ClNO_2) is
352 limited. Therefore, the production rate of ClNO_2 is generally low in areas affected by
353 fresh emissions. Since the contribution of direct emissions to Cl_2 is low and it is
354 predominantly produced secondarily in the atmosphere, high levels of Cl_2 are also
355 found away from the fresh emissions.

356 The spatial distribution of N_2O_5 concentrations differs from that of other species
357 (Figure 3c). While the concentrations of most of the species are higher in the southern
358 region, the N_2O_5 concentrations are lower in some parts of this region. This is because
359 the O_3 concentration in the core urban areas is low due to high NO_x emissions. The
360 N_2O_5 concentrations from the HET case are approximately 16% lower on average
361 (Figure 3d) because the Bertram and Thornton (2009) parameterization used in the HET
362 case generally gives higher uptake coefficients than the parameterization of Davis et al.
363 (2008) used in the BASE case (Table 3).

364 Although the higher uptake coefficients of N_2O_5 in the HET case facilitate faster
365 conversion of N_2O_5 to nitrate, the nitrate concentrations do not always increase. During
366 daytime hours, nitrate concentrations in the HET case increase due to higher OH (Figure
367 3e and Figure 3f). At nighttime, in contrast, the nitrate concentration decreases
368 significantly in some regions by about 22%, mainly due to lower molar yield of nitrate
369 from the N_2O_5 heterogeneous reaction in the HET case (Figure 3g and Figure 3h).
370 Although ClNO_2 produced in the N_2O_5 reaction also produces nitrate through a
371 heterogeneous reaction when the particle pH is above 2, which is true for most regions
372 (see Figure S4), the uptake coefficient of ClNO_2 is significantly lower than that of N_2O_5

373 (0.01~0.09 for N_2O_5 and 6×10^{-3} for ClNO_2), leading to an overall decrease of nitrate
374 production. As the ClNO_2 production from the heterogeneous reaction leads to less
375 N_2O_5 conversion to non-relative nitrate, it may change the overall lifetime of NO_x and
376 their transport distances. The magnitude of this change and its implications on ozone
377 and $\text{PM}_{2.5}$ in local and downwind areas should be further studied.

378

379 3.4 Relationship between nitrate formation and chlorine chemistry

380 Nitrate productions from the homogeneous and heterogeneous pathways in Beijing are
381 approximated by the difference in predicted nitrate concentrations between the BASE
382 or HET case and a sensitivity case without heterogeneous reactions. Averaging over the
383 five-day period, approximately 58% of the nitrate originates from HNO_3 gas-to-particle
384 partitioning and 42% is from heterogeneous reactions (Figure 4). This conclusion
385 generally agrees with measurements at Peking University (PKU) (52% from the
386 heterogeneous process and 48% from HNO_3 partitioning) on four polluted days
387 (average in September 2016 reported by Wang et al. (2017)). Slightly higher
388 contributions of the homogeneous pathway in this study is expected because of high
389 OH concentrations during the day and lower particle surface areas at night.

390 The nitrate formation from different homogeneous and heterogeneous pathways in
391 the BASE case and HET case are further studied. Contributions of different gas phase
392 pathways are determined using the process analysis tool in CMAQ. Contributions of
393 different heterogeneous pathways are determined using a zero-out method that turns off
394 one heterogeneous pathway at a time in a series of sensitivity simulations. Figure 4
395 shows that the reaction of OH and NO_2 is always the major pathway for the formation
396 of nitrate through homogeneous formation of HNO_3 and gas-to-particle partitioning.
397 However, its nitrate production rate through this homogeneous pathway decreases
398 significantly from daytime to nighttime (from $1.81 \mu\text{g m}^{-3} \text{h}^{-1}$ to $0.33 \mu\text{g m}^{-3} \text{h}^{-1}$ on
399 average). The nitrate production from other HNO_3 partitioning pathways in the daytime
400 is negligible. At nighttime, homogeneous reaction of N_2O_5 with water vapor accounts
401 for approximately 5% of the overall homogeneous nitrate formation. For the
402 heterogeneous pathways, daytime production rate is approximately $0.6 \mu\text{g m}^{-3} \text{h}^{-1}$ with

403 1/3 of the contributions from NO₂ and 2/3 from N₂O₅. Nighttime production on nitrate
404 from the heterogeneous pathways is approximately 3.1 μg m⁻³ h⁻¹, of which 85% is due
405 to N₂O₅ and 15% is due to NO₂.

406 Comparing the BASE case and the HET case shows that, when the chlorine
407 chemistry is included, the gaseous HNO₃ produced by OH reacting with NO₂ increases
408 significantly in the HET case. Correspondingly, the nitrate production rate reaches up
409 to 2.04 μg m⁻³ h⁻¹ in the daytime due to increased atmospheric OH concentrations
410 predicted by the chlorine reactions. Similar conclusions are also obtained by Li et al.
411 (2016) and Liu et al. (2017) based on observations and model simulations. The
412 heterogeneous production of nitrate from the reaction of N₂O₅ uptake decreases by
413 approximately 27% in the HET case due to the production of gas phase ClNO₂.
414 According to the study by Sarwar et al. (2012; 2014), including the heterogeneous
415 reaction of N₂O₅ with PCl decreased the nocturnal nitrate concentration by 11-21% in
416 the United States, which was slightly less than the current study for Beijing. It is likely
417 because PCl concentrations in the United States are significantly lower than those in
418 Beijing (the monthly PCl concentration is 0.06 μg m⁻³ in the United State against ~1 μg
419 m⁻³ in Beijing) so that PCl is depleted quickly. The contributions of NO₂ uptake to
420 nitrate also decrease by 22% because of the lower rate constant of the reaction of NO₂
421 with PCl. In contrast, the contribution of ClNO₂ reacts with particle surface to nitrate
422 production increases by 6% in the HET case. The overall nitrate concentration in the
423 HET case is about 22% higher than that in the BASE case during this study period.

424

425 **4. Conclusions**

426 In this work, a modified CMAQ model incorporated with heterogeneous reactions for
427 the production of molecular chlorine and other reactive chlorine species is developed
428 and its impact on of the nitrate formation predictions are evaluated. The contributions
429 from different homogenous and heterogeneous pathways to nitrate formation are also
430 quantified. High concentration of Cl₂ and ClNO₂ do not occur in the southern part of
431 the Beijing-Tianjin-Hebei region with intensive emissions of chlorine species as higher
432 concentrations of O₃ and N₂O₅ associated with the heterogeneous formation of these

433 species generally occurred in the downwind areas. CTMs without a complete treatment
434 of the chlorine chemistry can underestimate daytime nitrate formation from the
435 homogeneous pathways, particularly from HNO₃ gas-to-particle partitioning due to
436 underestimation of OH concentrations and overestimate the nighttime nitrate formation
437 from the heterogeneous pathways due to missing chlorine heterogeneous chemistry.

438

439 **Data availability.** The data in this study are available from the authors upon request
440 (shxwang@tsinghua.edu.cn)

441

442 **Author contributions.** XQ, QY, SW and JH designed the study; YS, BL, AS, XY
443 provided observation data; XQ, QY, SW, JZ, QX, DD, LD and JX analyzed data. XQ,
444 QY and SW wrote the paper.

445

446 **Competing interests.** The authors declare that they have no conflict of interest.

447

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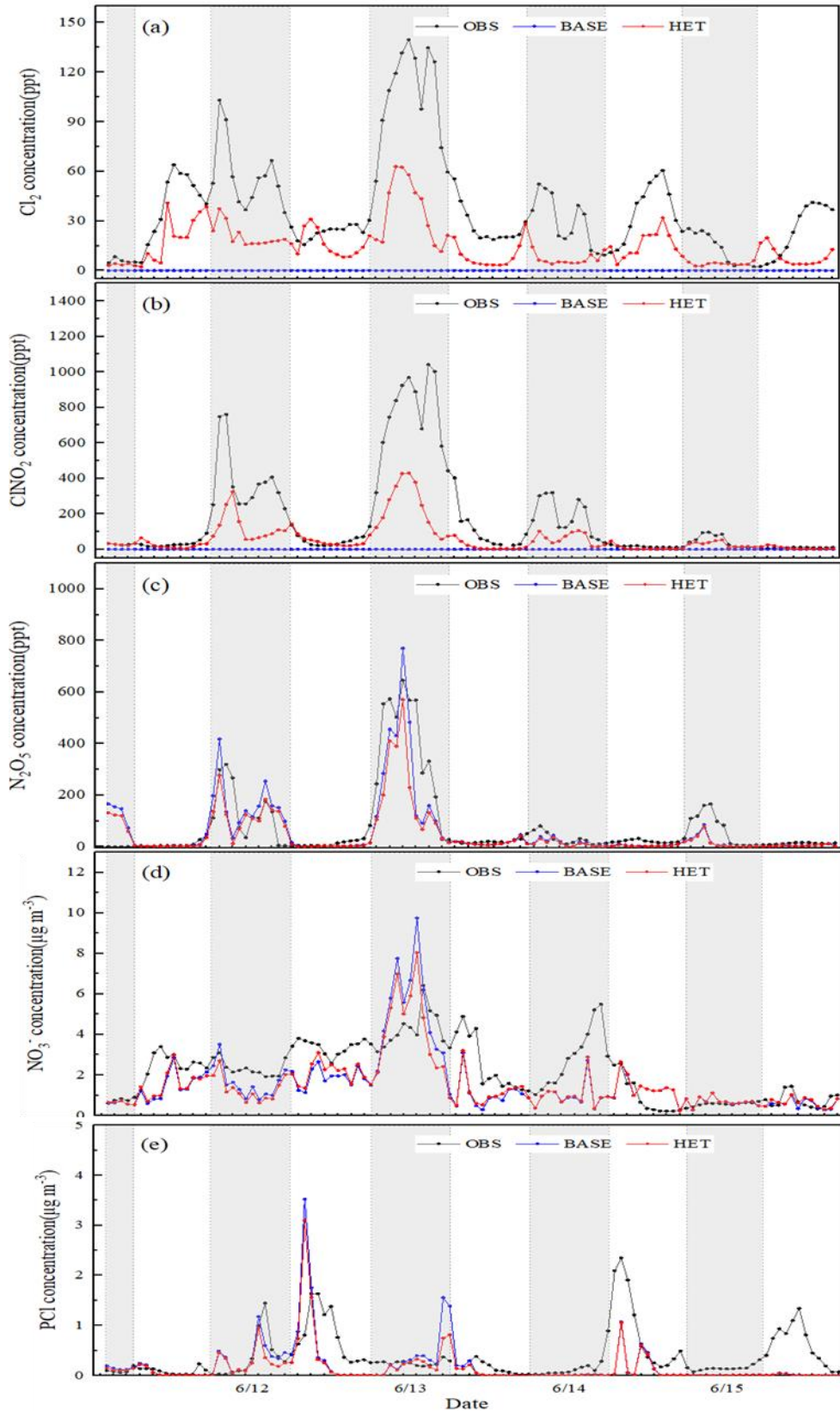
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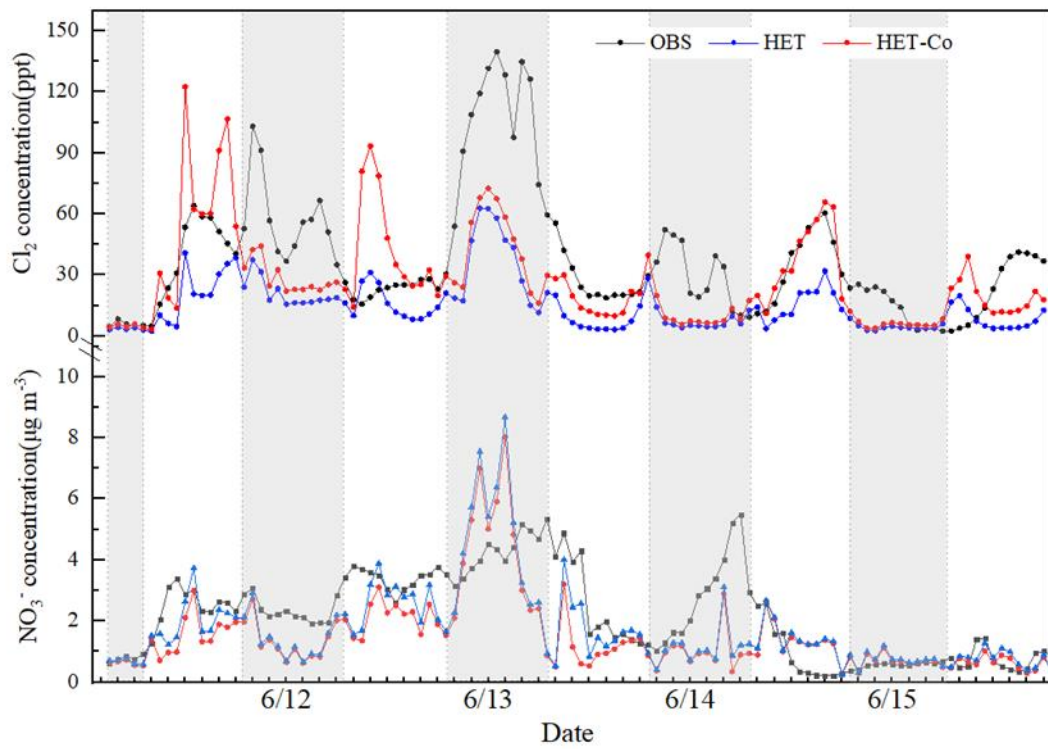
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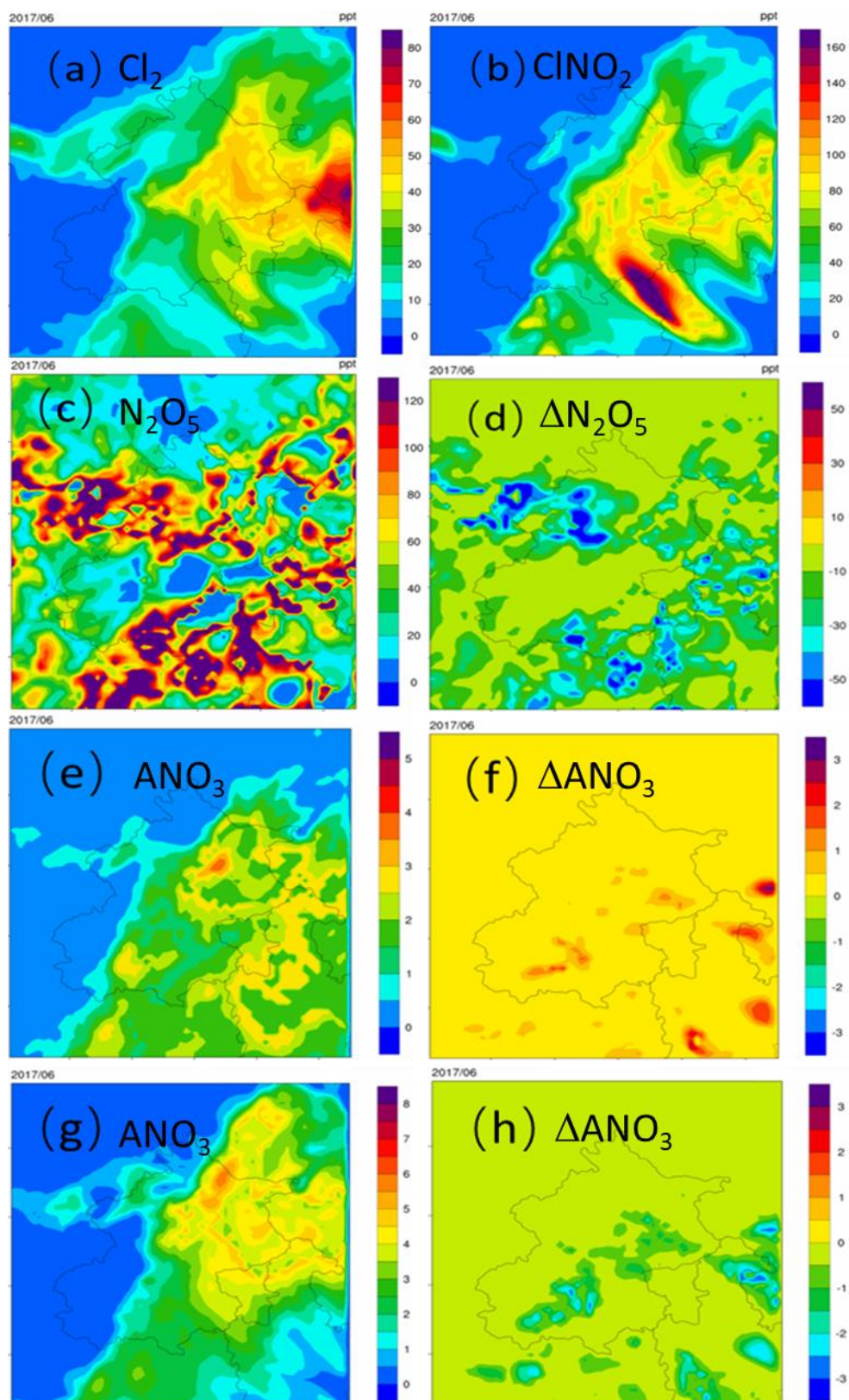
590 Figure 1 Comparison of observed hourly Cl_2 , ClNO_2 , N_2O_5 (at the Institute of
 591 Atmospheric Physics, Chinese Academy of Sciences), NO_3^- and PCl (at Tsinghua
 592 University) in urban Beijing with predictions from the BASE and the HET cases
 593 during 11-15 June 2017.



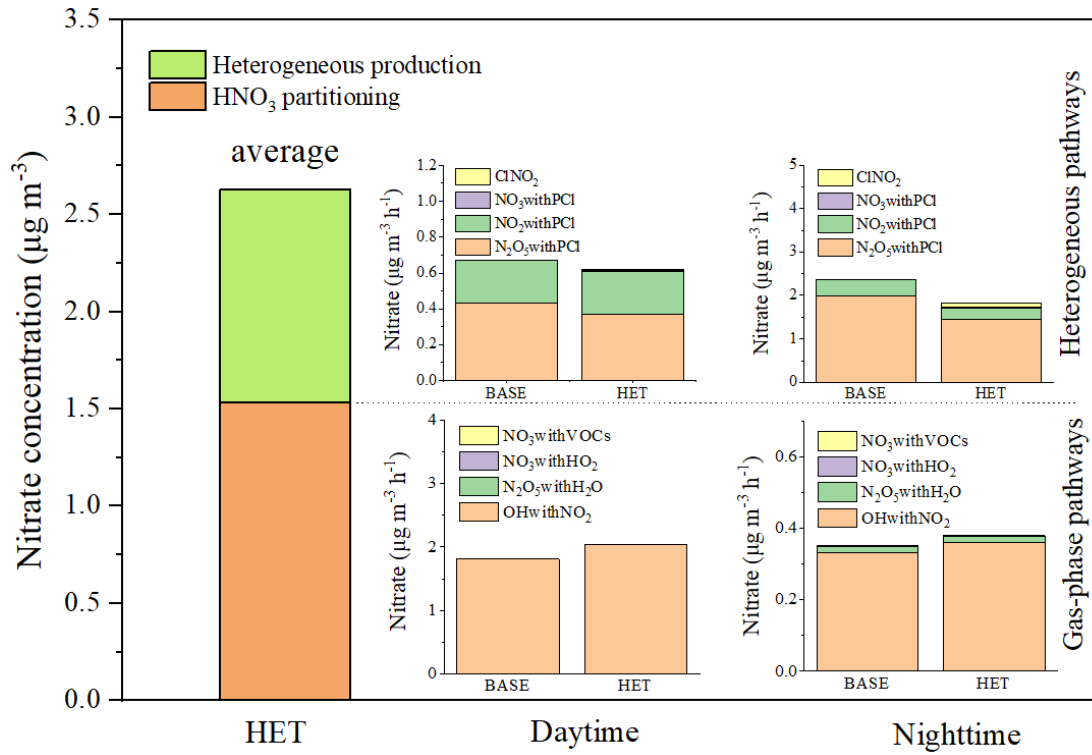
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595 Figure 2 Comparison of observed and predicted Cl_2 and NO_3^- concentrations under
 596 different uptake coefficient of O_3 (HET: daytime $\gamma_{\text{O}_3} = 1 \times 10^{-3}$, nighttime $\gamma_{\text{O}_3} =$
 597 1×10^{-5} ; HET-Co: daytime $\gamma_{\text{O}_3} = 1 \times 10^{-2}$, nighttime $\gamma_{\text{O}_3} = 1 \times 10^{-4}$).

598



599
 600 Figure 3 Spatial distributions of episode-average (a) Cl_2 , (b) ClNO_2 , (c) N_2O_5 , (e)
 601 daytime (8:00 – 20:00 hours) nitrate (ANO_3) and (g) nighttime nitrate concentrations
 602 from 11-15 June 2017, and the differences in the episode-average (d) N_2O_5 (HET case
 603 – BASE case), (f) daytime nitrate and (g) nighttime nitrate. Units are $\mu\text{g m}^{-3}$.
 604



605

606 Figure 4 Contributions of different homogeneous and heterogeneous pathways to nitrate
 607 formation.

608

609 Table 1 The sectoral emissions of HCl, Cl₂ and PCl in Beijing in 2017. Unit: Mg year⁻¹
 610 ¹

Sector	Emissions		
	HCl	Cl ₂	PCl
Power plant	22.8	1.2	6.75
Industry	587.3	20.1	89.2
Residential	202.4	8.1	34.7
Biomass burning	0.182	0	0.14
Municipal solid waste	1080.2	0	8.47
Cooking	0	0	426.8
Total	1892.9	29.4	566.1

611

612

613 Table 2 Major gas-phase and heterogeneous pathway of producing nitrate in original CMAQ and newly added or revised heterogeneous reactions
 614 in improved CMAQ.

Type	Reactions	No.	Reference	Comment
Original CMAQ				
Gas-phase chemistry	$\text{OH} + \text{NO}_2 \rightarrow \text{HNO}_3$	R1		
	$\text{N}_2\text{O}_5 + \text{H}_2\text{O} \rightarrow 2\text{HNO}_3$	R7		
	$\text{HO}_2 + \text{NO}_3 \rightarrow 0.2\text{HNO}_3 + 0.8\text{OH} + 0.8\text{NO}_2$	R8		
	$\text{NO}_3 + \text{VOCs}^a \rightarrow \text{HNO}_3$	R9		
Heterogeneous chemistry	$\text{N}_2\text{O}_5(\text{g}) + \text{H}_2\text{O}(\text{aq}) \rightarrow 2\text{H}^+ + 2\text{NO}_3^-$	R5		
	$2\text{NO}_2(\text{g}) + \text{H}_2\text{O}(\text{aq}) \rightarrow \text{HONO}(\text{g}) + \text{H}^+ + \text{NO}_3^-$	R10		
Improved CMAQ				
Newly added or revised heterogeneous reactions	$\text{N}_2\text{O}_5(\text{g}) + \text{H}_2\text{O}(\text{aq}) + \text{Cl}^-(\text{aq}) \rightarrow \text{ClNO}_2(\text{g}) + \text{NO}_3^-$	R6	Bertram and Thornton (2009)	Revise R5
	$2\text{NO}_2(\text{g}) + \text{Cl}^- \rightarrow \text{ClNO}(\text{g}) + \text{NO}_3^-$	R11	Abbatt et al. (1998)	Revise R10
	$\text{NO}_3(\text{g}) + 2\text{Cl}^- \rightarrow \text{Cl}_2(\text{g}) + \text{NO}_3^-$	R12	Rudich et al. (1996)	Increase NO_3^-
	$\text{O}_3(\text{g}) + 2\text{Cl}^- + \text{H}_2\text{O}(\text{aq}) \rightarrow \text{Cl}_2(\text{g}) + \text{O}_2(\text{g}) + 2\text{OH}^-$	R13	Abbatt et al. (1998)	Affect OH
	$2\text{OH}(\text{g}) + 2\text{Cl}^- \rightarrow \text{Cl}_2(\text{g}) + 2\text{OH}^-$	R14	George et al. (2010)	Affect OH
	$\text{ClONO}_2(\text{g}) + \text{Cl}^- \rightarrow \text{Cl}_2(\text{g}) + \text{NO}_3^-$	R15	Deiber et al. (2004)	Affect OH
	$\text{HOCl}(\text{g}) + \text{Cl}^- + \text{H}^+ \rightarrow \text{Cl}_2(\text{g}) + \text{H}_2\text{O}$	R16	Pratte et al. (2006)	Affect OH
	$\text{ClNO}_2(\text{g}) + \text{Cl}^- + \text{H}^+ \rightarrow \text{Cl}_2(\text{g}) + \text{HONO}(\text{g})$ (pH < 2.0)	R17	Riedel et al. (2012)	Affect OH
	$\text{ClNO}_2(\text{g}) + \text{H}_2\text{O}(\text{aq}) \rightarrow \text{Cl}^- + \text{NO}_3^- + 2\text{H}^+$ (pH ≥ 2.0)	R18	Rossi (2003)	Increase NO_3^-

615 ^a: presents different VOCs species. In the SAPRC-11 mechanism, the VOCs species include CCHO (Acetaldehyde), RCHO (Lumped C3+
 616 Aldehydes), GLY (Glyoxal), MGLY (Methyl Glyoxal), PHEN (phenols), BALD (Aromatic aldehydes), MACR (Methacrolein), IPRD (Lumped
 617 isoprene product species).

618

Table 3 Observed day (D) and night (N) NO_3^- concentrations (Obs.) and predicted uptake coefficient of N_2O_5 ($\gamma_{\text{N}_2\text{O}_5}$) and nitrate concentrations (Pred.) using the parameterizations of $\gamma_{\text{N}_2\text{O}_5}$ by Bertram and Thornton (2009) (Scenario 1), Davis et al., (2008) (Scenario 2) and the upper-limit value derived by Zhou et al. (2018) (Scenario 3)

	NO_3^- Obs.	Scenario1		Scenario2		Scenario3	
		$\gamma_{\text{N}_2\text{O}_5}$	Pred.	$\gamma_{\text{N}_2\text{O}_5}$	Pred.	$\gamma_{\text{N}_2\text{O}_5}$	Pred.
06/11-D	2.54	0.033	1.59	0.008	1.32	0.09	2.17
06/11-12-N	2.42	0.043	1.67	0.037	1.37	0.09	2.12
06/12-D	3.39	0.028	2.16	0.032	2.74	0.09	3.13
06/12-13-N	4.24	0.021	4.02	0.022	4.05	0.09	6.04
06/13-D	2.57	0.012	1.18	0.008	1.06	0.09	2.47
06/13-14-N	4.10	0.022	4.45	0.022	4.45	0.09	7.13
06/14-D	0.95	0.001	1.34	0.001	1.33	0.09	1.64
06/14-15-N	2.75	0.013	1.00	0.007	0.96	0.09	2.33
06/15-D	0.75	0.001	0.66	0.001	0.66	0.09	1.11