The file includes

- 1. Responses to the reviewer#1' comments
- 2. Responses to the reviewer#2' comments
- 3. The revised manuscript

General Response: We thank the reviewer for your helpful comments. We have addressed all comments and provided point by point response below. The revised manuscript is presented in below.

Response to the referee#1' comments

Yu et al. present findings from detailed compositional measurements of Arctic aerosol in Svalbard during August 2012. While there is obvious importance of conducting detailed physiochemical characterizations of Arctic aerosol in terms of their radiative impacts and subsequent indirect effects on frozen surfaces, there are major issues with the manuscript by Yu et al. that would need to be addressed prior to publication. These issues stem from possible misinterpretation of the data that shape the reported main findings. It would behoove the authors to provide a sufficient level of detail on the methodologies (including caveats) and results to support the main conclusions they report.

Response: We carefully addressed all of the questions and concerns raised.

General comments:

There is a scarcity of detail regarding which samples and particles were analyzed. More specifically, which samples were analyzed, which particles were analyzed per sample and how those were chosen, how many particles per sample were analyzed, and why only select samples and particle numbers were analyzed under each method is not at all defined.

Response: We added more details in the Experimental section. In addition, we revised Table S1 and added information on what samples were analyzed by what methods.

which samples were analyzed?

Response: Revised in the context line 144-146

"The sample information such as local sampling date and time and meteorological conditions (e.g., temperature (T), relative humidity (RH), pressure (P), wind direction (WD), and wind speed (WS)) are listed in Table S1."

which particles were analyzed per sample and how those were chosen,

Response: We follow a commonly used methodology in single particle analysis community. Since the distribution of aerosol particles on TEM grids was not uniform, with coarser particles occurring near the center and finer particles on the periphery five areas were chosen from the center and periphery of the sampling spot on each grid to ensure that the analyzed particles are representative of the whole sample.

We added the following in the revised manuscript. Line 159-162

"The distribution of aerosol particles on TEM grids was not uniform, with coarser particles occurring near the center and finer particles on the periphery. Therefore, to ensure that the analyzed particles are representative, five areas were chosen from the center and periphery of the sampling spot on each grid. Through a labor-intensive operation, 2002 aerosol particles with diameter < 10 μ m in 21 samples were analyzed by TEM/EDS (Table S1)."

How many particles per sample were analyzed, For example, which of the samples constituted the 2002 and 575 particles analyzed by TEM and TEM/EDS, respectively?

Response: Through a labor-intensive operation, 2002 aerosol particles with diameter < 10 μ m in 21 samples were analyzed by TEM/EDS (Table S1). To check composition of individual particles, EDX was manually used to obtain EDS spectra of individual particles. In the clean Arctic air, there are several relatively easy-to-identify particle types including sea salt, sulfate, soot, and OM. Because soot particles have chain-like aggregation, it is not necessary to check their elemental composition. Sea salt particles display spherical or square shapes and are stable under the electron beam in TEM but sulfate particles are spherical but flats on the substrate and produce unstable bubble under the electron beam (Buseck and Posfai, 1999; Chi et al., 2015). TEM observations also can clearly identify sulfate particles or sulfate with OM coating. Therefore, we can identify Arctic particle types based on their morphology. We usually randomly chose 20-30 particles in each sample for elemental analysis to confirm the identification of particle types (Table S1). In total, EDS spectra of 575 particles were manually obtained and saved in the computer for elemental composition analysis. Detailed information is now added to Table S1 We would like to point out that it is not realistic to analyse every single particle collected on the grid as each EDS analysis took about 100 s and all data need to be analysed manually. Therefore, we In the revised manuscript (track changed), we added the following (line 165 to 180)

"In the clean Arctic air, there are simply particle types including sea salt, sulfate, soot, and OM. Because soot particles have chain-like aggregation, it is not necessary to check their elemental composition. Sea salt particles display spherical or square shapes and are stable under the electron beam in TEM but sulfate particles are spherical but flats on the substrate and produce unstable bubble under the electron beam (Buseck and Posfai, 1999; Chi et al., 2015). TEM observations also can clearly identify sulfate particles or sulfate with OM coating. Therefore, we can easily identify Arctic particle types based on their morphology. Because of the time-consuming in the experiment, it is not necessary to frequently check elemental composition of 20-30 particles in each sample (Table S1). EDS spectra of 575 particles examined by TEM were dry at the time of observation in the vacuum of the electron microscope. In our study, the effects of water and other semi-volatile organics were not considered as they evaporate in the vacuum."

For certain techniques, only a few samples (i.e., 3 samples for NanoSIMS but no mention of particle number) or even only a handful of particles (i.e., only 17 particles for AFM but no mention of which sample(s) these came from) were analyzed, and in the case of SEM there is no information on sample or particle number. I understand that some of these tools, i.e., AFM, are time-consuming which is why a low number of particles were analyzed, but then the authors need to be careful about overstating result interpretations. It is important to know how many particles and from which samples to provide sufficient statistics and afford information on daily

source variability. As it stands, there is no way to tell how representative the percentages (which are hidden in the text) are of summertime aerosol in general, or just of specific samples from select days.

Response: Information is given in Table S1.

For some methods, we analysed only a small number of particles. This does not affect our conclusion. The purpose of AFM is to calibrate the equivalent circle diameter to equivalent spherical diameter. As the previous studies (Chi et al., ACP, 2015), the number of samples analysed is enough to address this issue.

The purpose of NanoSIMS is to confirm the OM coating. we analysed 32 S-OM particles, which have the same morphology and composition; analyzing more samples is unlikely adding more information.

Table S1
* number of particles analysed

Date	Local time	Т	RH	Р	WD	WS	TEM	EDX	SEM	AFM	NanoSIMS
2012.8.7	20:50 -21:15	4.9	84	1009.0	296	4.1	43	10			
2012.8.8	08:23 -08:48	4.9	81	1007.6	238	2.1	38	11			
2012.8.9	14:40 -15:05	6.6	81	1003.9	129	6.5	146	50			12
	15:20 -15:49	7.0	78	1003.5	120	7.3	130	26	20		
2012.8.10	00:15 -00:40	7.3	80	998.6	135	8.9	121	23			
2012.8.11	09:10 -09:35	6.2	94	997.0	303	3.3	128	50			10
2012.8.11	16:00 -16:25	4.1	92	1002.0	327	4.6	156	55		6	
2012.8.12	15:25 -15:50	5.7	83	1006.8	132	6.9	100	15	32		
2012.8.13	08:55 -09:20	5.3	81	1009.6	91	1.1	113	16			
2012.8.13	14:15 -14:40	4.5	90	1011.4	351	2.1	136	56			10
2012.8.14	09:50 -10:20	5.0	85	1019.7	351	2.3	134	24			
2012.8.14	15:12 -15:42	4.6	88	1020.5	117	2.6	121	26			
2012.8.14	21:17 -21:47	4.8	84	1020.7	276	5.4	178	56		5	
2012.8.15	09:15 -09:45	5.8	73	1019.6	135	3.7	165	60		6	
2012.8.15	15:00 -15:33	6.8	70	1018.9	270	3.3	80	11			
2012.8.17	9:00 -10:00	3.8	86	1017.1	116	0.3	30	15			
2012.8.17	14:50 -15:20	3.7	85	1015.7	109	2.2	42	16			
2012.8.21	15:05 -15:40	1.6	87	1003.7	314	6.8	46	18			
2012.8.22	08:55 -09:30	2.8	78	999.2	331	2.8	49	19			
2012.8.23	09:00 -09:40	3.4	64	998.0	136	6.9	21	9			
2012.8.23	20:35 -21:08	3.8	59	1002.0	138	6.3	25	9			

It would be helpful to provide a figure or two of the overall picture of aerosol composition, e.g., bar graphs or pie charts.

Response: Bulk aerosols were not determined in this study. We provide a summary of elemental compositions of individual particles (see figure below and Supplementary Fig. 3). These are based on elemental compositions of EDS. Based on the Figure S3, O, Na, S, Cl are most abundant elements in the arctic particles.

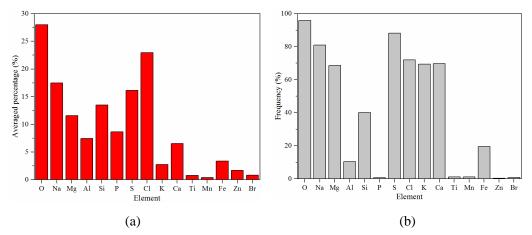
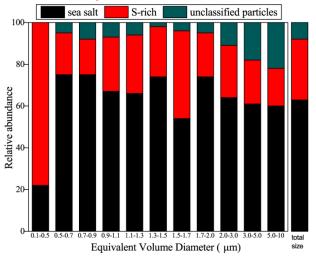
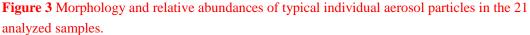


Figure S3 Elemental compositions of individual particles from EDS spectra. Left: Average weight of elemental compositions derived from the EDS spectra (b) frequency of element occurring in individual particles.

There are percentages provided in the text, but showing the relative abundance of each particle type is pretty standard. Along relative abundance, the authors report that 29% of the particles were non-sea salt. What percentage were unclassified? What percentage is the "majority of NSS-particles"?

Response: Indeed, there are a few of unclassified particles. 63% of particles were identified as the sea salt particles, 29% particles were NSS-sulfate particles, and 10% particles were unclassified particles. The Figure as the referee requests was added in the main manuscript.





As a result, it is not clear how important these particles are in general in the context of radiative

impacts, given sea salt was what seemed like the dominant particle type and also largely affects scattering and SSA. In addition, to demonstrate that these particle types are important for the Arctic energy budget, they should show extinction properties for total aerosol (including sea salt) in Figure 6. I would think given the typical sizes for these types of particles (sulfate and soot) and reported abundance for this particular study, they would not affect the scattering cross sections relative to sea salt. The emphasis on the radiative impacts of these aerosol types is a large part of the manuscript, so their properties need to be presented in broader context. Are they important within the total aerosol population or not? This would be more relevant to the actual atmospheric implications.

Response: As suggested by the reviewer, the scattering cross section of sea salt particles is much higher than other aerosols. On the other hand, the purpose of this study is to show the potential role of BrC coating on sulfate particles on absorption. since sea salts do not absorb light, it is the soot and BrC that matters for the light absorption in the atmosphere within the whole aerosol population. does, and so does BrC. In the Arctic, it is the soot and BrC that has the potential to warm up the climate, even though their total absorption is likely to significantly lower than the absolute value of scattering.

The sizes of sulfate and soot particles are given in Figure 8. Our analysis cannot provide a clear information on the radiation balance but it did suggest the potential role BrC, which we know plays an important role in Arctic climate

In the revised manuscript, we revised Figure 8 and added the following (line 472 to 482)

"Figure 8c also shows that the single scattering albedos (SSAs) of individual particles are 0.92, 0.99, and 1 when assuming the BrC as strongly, moderately and non-absorbing (cases SSA1 to SSA3). These results suggest whether we consider organic coating as BrC may have a significant influence on the absorption properties of individual sulfate particles.

In this study, we expored the relationship between ACS of individual particles and particle diameters. Interestingly, Figure 8d shows that ACS of individual fine OM-coating sulfate particles increased following the increasing particle size. The result shows that the ACS can be enhanced following particle size growing and particle aging. In other word, OM-coating sulfate particles transported more longer distances and they might have stronger optical absorption in the Arctic air."

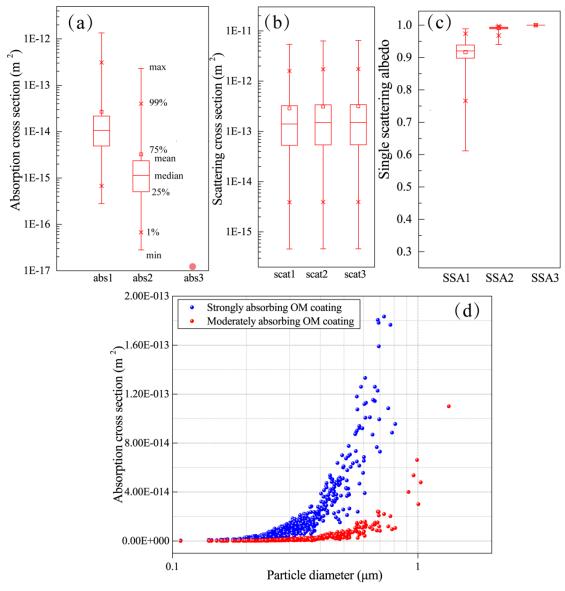


Figure 8 Optical properties of Box-and-whisker plots showing optical parameters of all analysed particles assuming sulfate core and BrC shell (not considering soot cores in the particles). (a) Scattering cross section (b) Absorption cross section (c) Single scattering albedo. Top to bottom makers in the box-and-whisker represent max, 99%, 75%, mean, median, 25%, 1%, min values. (d) Absorption cross section along with particle diameter assuming strongly absorbing BrC and Moderate absorbing BrC as the particle OM coating.

There are several issues with the methods as presented. For instance, there is very little detail given on how the particles were classified under each technique, there are no errors or statistical analyses reported, and certain methods have very little specification detail (i.e., SEM is a very short paragraph). Regarding the samples storage, at 20% RH, I would assume all volatile and most semi-volatile species would evaporate, significantly altering particle shape, size, and composition. I will admit, these techniques are not my area of expertise, but the authors should at least comment on potential losses and caveats with this storage method. If there are significant losses of material, how representative are the analyzed particles of the total ambient aerosol population at the time of collection? I am skeptical the authors are comparing apples to apples

by possible alteration of particles during storage.

Response: We add more specification detail of SEM in the paragraph.

SEM/ TEM/NanoSIMS all have to be analysed under vacuum so the volatile fractions will be lost during the analyses, but there is absolutely no other way to observe the shape of the particles from the ambient air. Laskina et al., EST, (2015) tested different methods and confirmed that the best way is settled in dry condition for further TEM or SEM analysis. Individual particle analyses offer the advantages of the mixing state and composition of individual particles.

The details about individual particle analysis have been reviewed by one previous paper (Li et al., JCP, 2016)

Laskina, O., Morris, H.S., Grandquist, J.R., Estillore, A.D., Stone, E.A., Grassian, V.H., Tivanski, A.V., 2015. Substrate-Deposited Sea Spray Aerosol Particles: Influence of Analytical Method, Substrate, and Storage Conditions on Particle Size, Phase, and Morphology. Environ. Sci. Tech. 49 (22), 13447-13453.

Li, W., Shao, L., Zhang, D., Ro, C.-U., Hu, M., Bi, X., Geng, H., Matsuki, A., Niu, H., Chen, J., 2016. A review of single aerosol particle studies in the atmosphere of East Asia: morphology, mixing state, source, and heterogeneous reactions. J. Clean. Prod. 112, Part 2, 1330-1349.

We've added the following the revised manuscript: L141-144

"Ambient laboratory conditions (17–23% RH and 19–21 °C) is effective at preserving individual hygroscopic aerosol particles and reducing changes that would alter samples and subsequent data interpretation (Laskina et al., 2015)."

L178-180"Particles examined by TEM were dry at the time of observation in the vacuum of the electron microscope. In our study, the effects of water and other semi-volatile organics were not considered as they evaporate in the vacuum."

Regarding the source analysis, there is very little detail given on the FLEXPART modeling and only a couple sentences on the results and discussion of the simulations. It is used to a very minimal extent and very generally summarized, even though Arctic aerosol sources can vary drastically day-by-day, and especially given possible local contributions. Figure 5 is very difficult to discern and glean any source information from it. Also, why are only these particular days shown?

Response: We revised the part related to the FLEXPART and added the sources around Arctic areas. Here we also added the Figure 1 that showing back trajectory of air mass during the sampling periods.

The reason we shosed the FLEPART at the specific days is that we did FLEPART based on the preliminary works including back trajectories of each sampling and TEM study. During these days, We found abundant sulfate and soot in samples in 9-15, Aug (Table S1). This is the reason that we planned to the FLEXPART modeling.

Figure 2 could not give direct source information but they can provide potential source locations. Here we added two Figures in the supplemental which provide the emission intensity in the Arctic area.

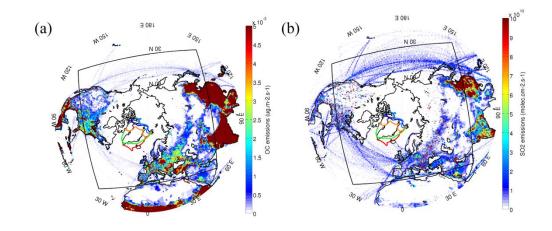


Figure S2 OC (a) and SO₂ (b) Emission intensity in Arctic area and 24h back trajectories on August 11, 12, 14, and 15, 2012.

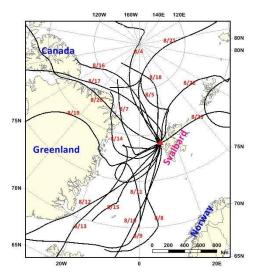


Figure 1 72 h back trajectories of air masses at 500m over Arctic Yellow River Station in Svalbard during 3–26 August 2012, and arriving time was set according to the sampling time

There is no background on previous relevant studies conducted at the study location, even though there is a long-term monitoring station with aerosol measurements at Ny-Ålesund (https://www.esrl.noaa.gov/psd/iasoa/stations/nyalesund). It is not the same exact location as the Chinese site, but close enough to at least use those routine, publically-available measurements to provide some broader spatial and temporal context.

Response: this is now being revised. We've added L127-130

"The sampling site is about 2 km far away from the Zeppelin observatory station (78.9N 11.88E) running by the Ny-Ålesund Science Managers Committee (https://www.esrl.noaa.gov/psd/iasoa/stations/nyalesund). Two to three samples were regularly collected at 9:00, 16:00, 21:00 (local time) of each day, with a total of 46 samples during 7-23 August, 2012."

There is only basic mentioning of biogenic VOCs, but none of biogenic or biological aerosol. The Arctic summer, especially in remote coastal sites, is largely affected by gases and aerosol from primary productivity due to the availability of sunlight and open water. There is no discussion on if the OM is biogenic/biological, and in general, the definition of OM is vague.

Response: The methodology we use cannot identify the sources of the SOAs, whether biogenic or non-biogenic. In the arctic atmosphere, Leck and Svensson (2015) found some biogenic aerosols like gel-aggregate containing bacterium in ultrafine particles. In our study, we didn't collect ultrafine particles using the sampler.

We've added the following to line L405-408

"Similarly, besides the OM coating in the Arctic particles, Leck and Svensson (2015) found some biogenic aerosols like gel-aggregate containing bacterium in ultrafine particles. However, we didn't find any gel-like particles in the samples because our sampler had very low efficiency for ultrafine particles."

Technical corrections:

There are many typos, grammatical issues, and a lack of necessary explanation (e.g., the 3 sets of bars in Figure 6, SEM a short paragraph with no numbers, etc.). Response: Thanks. We revised them.

References:

- Buseck, P.R., Posfai, M.: Airborne minerals and related aerosol particles: Effects on climate and the environment, P. Natl. Acad. Sci. USA, 96 (7), 3372-3379, 1999.
- Chi, J.W., Li, W.J., Zhang, D.Z., Zhang, J.C., Lin, Y.T., Shen, X.J., Sun, J.Y., Chen, J.M., Zhang, X.Y., Zhang, Y.M., Wang, W.X.: Sea salt aerosols as a reactive surface for inorganic and organic acidic gases in the Arctic troposphere, Atmos. Chem. Phys., 15 (19), 11341-11353, 2015.
- Laskina, O., Morris, H.S., Grandquist, J.R., Estillore, A.D., Stone, E.A., Grassian, V.H., Tivanski, A.V.: Substrate-Deposited Sea Spray Aerosol Particles: Influence of Analytical Method, Substrate, and Storage Conditions on Particle Size, Phase, and Morphology, Environ. Sci. Tech., 49 (22), 13447-13453, 2015.
- Leck, C., Svensson, E.: Importance of aerosol composition and mixing state for cloud droplet activation over the Arctic pack ice in summer, Atmos. Chem. Phys., 15 (5), 2545-2568, 2015.
- Maahn, M., de Boer, G., Creamean, J.M., Feingold, G., McFarquhar, G.M., Wu, W., Mei, F.: The observed influence of local anthropogenic pollution on northern Alaskan cloud properties, Atmos. Chem. Phys., 17 (23), 14709-14726, 2017.

General Response: We thank the reviewer for your helpful comments. We have addressed all comments and provided point by point response below. The revised manuscript is presented in below.

Response to the referee#2' comments

Overall Comments: The authors use multiple measures to investigate the physical properties of non-sea salt aerosols collected in the Arctic. They qualitatively describe the mixing state of organic, soot, and sulfate aerosols. Although the core data set of this manuscript is informative, as is, the presentation and discussion are confusing and possibly misleading. No large scale descriptive statistics are provided, the methods and presentation are often unclear and/or redundant, the environmental context is lacking, and the optical conclusions are not well explained. This manuscript would benefit from additional analysis, clarification of methods, adding proper context to results, and careful grammar review.

Response: We carefully addressed all of the questions below. We attempted to re-write the manuscript. However, we are not able to address non-specific questions.

Unclear Sample Selection: In multiple sections of this paper the authors discuss a subset of their samples without explaining why they were chosen. This detracts from their discussion and leaves the reader questioning what is not being discussed. Anytime a subset of samples is chosen or used, an explanation needs to be provided explaining why these are the best samples for that specific analysis or investigation. A few example occurrences can be found in lines 275, 281, 118, 190.

Response: Thank you for good comments. We received some comments before and revised the manuscript several rounds. After that, we added Figure S1 showing what samples were analyzed by different instruments. We removed some number which confused for the readers and rephrased some sentences in the Mehods section.

To SEM and NanoSIMS analysis after TEM observations, we need to select better samples. These samples should keep good conditions of carbon film because some samples have broken carbon film during the sampling period. In the NanoSIMS and SEM analysis, these samples need to be sticked on the sample stage, they could not be recycled again.

Please see response to reviewer 1, comments 2-4.

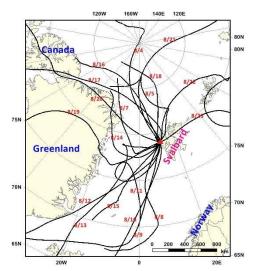
Environmental Context: The only environmental context explored in this text is the back trajectories provided for 4 of the sampling days. Whenever field data is discussed, external variables such as temperature, humidity, and back trajectories can help explain variability observed in the data. To improve this work, the context of the sampling should be explored as a driving factor between differences in particle characteristics.

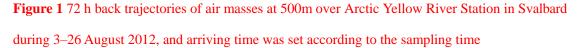
Response: We added the RH, T, P in the table S1. These meteorological data were recorded during the sampling period. Here we did back trajectories of each sampling day as shown in Figure 1.

We added the following to the revised manuscript (line 144 to 146)

"The sample information such as local sampling date and time and meteorological conditions

(e.g., temperature (T), relative humidity (RH), pressure (P), wind direction (WD), wind speed (WS)) were listed in Table S1."





2.1 I am also not convinced that the daily averaged back trajectory calculations are meaningful for these samples. (I'm also not sure the trajectories were daily averaged as the methods section just said they were calculated for a given day, not the frequency of times or given time on the day.) Because the sample length was between 20 minutes and 2 hours, a back trajectory occurring during the actual sampling time for all samples would be more appropriate. Response: We did choose the mid of the sampling time as the end point for the back trajectory analysis (see Figure 1)

3 Optical Property Calculations: The authors attempt to estimate the potential radiative forcing implications of their results using a core-shell model and Mie theory. Although this is a potentially meaningful result, the methodology needs to be discussed further and improved.

Response: We added more description in the Method section.

L263-271"BHCOAT Mie code by Bohren and Huffman (1983) was used to calculate the optical properties, including scattering cross section (SCS), absorption cross section (ACS), and single scattering albedo (SSA), assuming a core-shell structure. We firstly calculated these parameters assuming a sulfate core and OM shell structure only (ignoring some of the particles that contain soot core). Because the Mie code only can calculate the core-shell structure or homogeneous models, we assume sulfate as a core and OM as a shell in individual particle to build the core-shell model. Based on the core-shell standard mode (Li et al., 2016), we can calculate optical properties of individual internally mixed particles."

3.1 When performing Mie theory calculations the input size distribution characteristics and wavelength dependence (angstrom exponent) have dramatic effects on the results. These data

need to be reported. The authors state that a single size distribution is used. How was this size distribution established? Could you run your analysis over all of the size distributions you observed to estimate the variability?

Response: Size distributions of particles are measured by microscopes These measurements can help us to evaluate sulfate and OM volume and then we made the Figure 8d. Once we obtained these basic data, we can input all the data to calculate optical properties of individual particles. Here the result is different from the bulk online optical properties. As the comments, we did correction between optical absorption cross section (ACS) and particle diameter assuming strongly absorbing BrC and Moderate absorbing BrC as the particle OM coating. Interestingly, we found that ACS increase along with the particle diameter increase through one nonlinearity showed as below.

In this study, it is not necessary to consider wavelength dependence because lots of work how different size of aerosol particles interacted with the different wavelength. We just did pilot cases to confirm how BrC influence the optical absorption of sulfate particles based on our TEM data (Figure 8). The purpose of this study draws attention for BrC in the Arctic air.

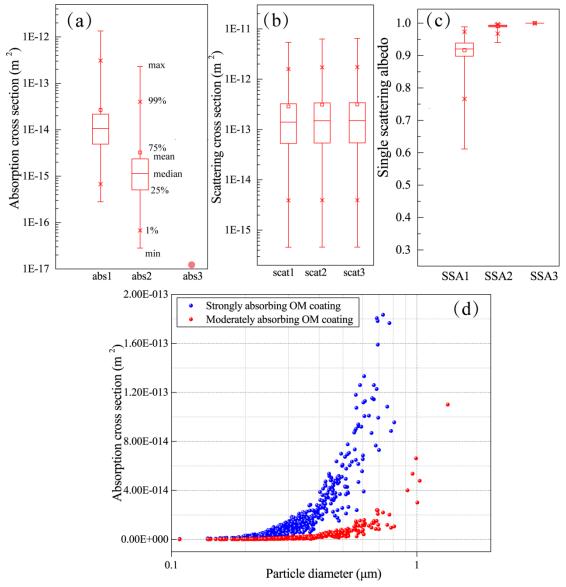


Figure 8 Optical properties of Box-and-whisker plots showing optical parameters of all

analysed particles assuming sulfate core and BrC shell (not considering soot cores in the

particles). (a) Scattering cross section (b) Absorption cross section (c) Single scattering albedo. Top to bottom makers in the box-and-whisker represent max, 99%, 75%, mean, median, 25%, 1%, min values. (d) Absorption cross section along with particle diameter assuming strongly absorbing BrC and Moderate absorbing BrC as the particle OM coating.

3.2 The way the authors include soot in their calculations is not appropriate, as soot is treated as a core with a mixed sulfate-OM shell. This is in direct conflict with their observations that soot is only observed to be associated with the OM shell and sulfate is always the core. It would be better not to include a soot calculation than include a misleading one.
Response: We deleted the soot part here. The updated data is shown in Figure 8

3.3 The relevancy of the refractive index (RI) choice needs to be discussed further. The authors use an RI from biomass burning brown carbon as their slightly absorbing case but frequently state that they believe the OM to be secondary. If that is the case, their estimations are likely too high for all cases but the non-absorbing case. The authors should include refractive indexes from secondary organic aerosol brown carbon to get a more realistic answer. These are generally much lower than the RI values that they used. Additionally, the scattering component chosen (1.65) is relatively high and not explained. Please explain why this value was chosen as it is especially important in core-shell cases.

Response: There are many different RIs from the laboratory experiments. For example, Jiang et al., (2019) suggested the measured RIs on different organic species. The RI is dependence on the wavelength. Fortunately, we just want to test how OM coating influence the sulfate particles. Based on the relationship between RIs and wavelength (Jiang et al., 2019; Feng et al., 2013), the RI doesn't influence our conclusions in this study.

Reference added: Jiang, H., A. L. Frie, A. Lavi, J. Y. Chen, H. Zhang, R. Bahreini, and Y.-H. Lin (2019), Brown Carbon Formation from Nighttime Chemistry of Unsaturated Heterocyclic Volatile Organic Compounds, Environmental Science & Technology Letters, DOI: 10.1021/acs.estlett.1029b00017.

Here we chose the RI reported in Feng et al., 2013. Through the comparisons between strongly absorbing and moderately absorbing, non-absorbing BrC, we generally knew how BrC coating could influence optical properties of sulfate particles. Although our current study could not get a more realistic answer, but the result suggests that BrC has a potentially role to play in light absorption properties of aerosols Arctic. The conclusion warrants further study in the Arctic area. Our results from the TEM can provide solid evidence about the mixing state of sulfate and other aerosol species. As suggested by the referee, we added a sentence to explain why we selected the 550 nm:

In the context:

"Although the refractive index has dependence on the wavelength between 350-870 nm, we tried to select the 550 nm as a case to test how OM coating influence sulfate particles in Arctic air."

3.4 Stemming from the above comment, the relevancy of each refractive index case needs to be discussed in the context of the Arctic. Which case seems the most likely?

Response: We can't make this judgement because there is no data to show this. We showed that this is potentially important to consider, which warrants further study.

3.5 The authors use m=1.55 for sulfate. RI values have wavelength dependence, was this included in any way?

Response: Thanks. It does. We corrected the writing here. Line 256 to 257: "The refractive index used for the non-light-absorbing sulfate component was set to m=1.55 at 550 nm (Seinfeld and Pandis, 2006)."

3.6 The authors average the absorption cross section on a per particle basis. This isn't meaningful since observed absorption cross sections will depend on the whole aerosol population (and the size of the particle). The authors should calculate an absorption cross section for the ambient aerosol concentration during their sampling and compare it to other absorption observations. Response: Our methods are not able to give the ambient aerosol concentration. We tested how the absorbing OM influenced optical properties of sulfate in the Arctic based our individual particle measurements. These results showed that this is potentially important to consider, which warrants further study.

4 Redundancy and Clarity: The writing of this paper needs improvement. There are multiple times when the authors restate the same point twice or fail to introduce a topic before discussing their results. An example of redundancy can be observed when Copper TEM grids are introduced in lines 120, 140, and 156. An example of an improper discussion occurs in the discussion of satellite particles, which are introduced in line 280 with no context or explanation. This forces the reader to infer what the authors mean by satellite particles, possibly leading them to wrong conclusions. To improve the manuscript, I recommend careful reconsideration of the presentation of the data, with special consideration to avoiding redundancy and ensuring the appropriate context is present.

Response: We carefully revised them. Please see our replies as below associated with your specific comments.

4.1 There are multiple points in the methods, a few of which I've included in the specific comments, that are unclear or confusing. The reader needs to be able to understand exactly which analysis was performed on which filters for how many particles if they are to believe your result.

Response: We made one major revision in this section. We added the Table S1 to make specific information.

5 Sample Information: The only summary of the total data set is provided in figure S6 and this figure states that only 3 samples (of 46 collected and 21 analyzed). No explanation is given as to why is summary is so limited given the authors have EDS data (which they used for classification) of 20-30 particles over 21 samples (at least 400 data points). All the of following discussions only make sense if they are provided in the context of overall sample composition.

Response: The Referee is probably referring to Figure S4 rather than Figure S6. It is a spelling

mistake, which we have corrected. At the revised ms, the Figure was moved to the main ms as Figure 1.

Specific Comments:

1 There are many grammatical errors and redundancies that I have not addressed below. Response: We carefully improved English writing in the manuscript.

2 The phrase internal mixing is used throughout the manuscript without an explicit definition. An explanation of what exactly you mean when you say something is "internally mixed" would improve the manuscript

Response: We added the following to the text Add an explanation the first time you write it, e.g., line 88 "Internal mixing means that a single particle simultaneously contains two or more types of aerosol components (Li et al., 2016)."

3 Line 55: Change "Artic" to "the Artic"Response: Corrected4 Line 60: Change "nature" to "natural"Response: Corrected

5 Line 64-66: Restructure this sentence for clarity. Treat the percentages in a consistent way as to not confuse the reader. For example, this sentence could be changed to:" For example, Winger et al.(2017) showed most Arctic BC is sourced from domestic activities (35%) and transportation (38%), with only minor contributions from gas flaring (6%), power plants (9%), and open fires (12%)" Response: Corrected

6 Line 72: I'm not sure exactly how this sentence fits in with the brown carbon theme of this paragraph. Are these compounds commonly found in brown carbon or organic aerosols in general? Please add some context.

Response: Thank you very much. We move the positions of this sentence and add more description here.

L73-78: "Accumulation of secondary organic aerosol, a significant fraction of the new particles grow to sizes that are active in cloud droplet formation in the Arctic (Abbatt et al., 2019). More than 100 organic species were detected in the Arctic aerosols and polyacids were found to be the most abundant compound class, followed by phthalates, aromatic acids, fatty acids, fatty alcohols, sugars/sugar alcohols, and n-alkanes (Fu et al., 2008)."

7 Line 84: replace "were" with "have been" Response: Corrected

8 Line 96-68: This sentence is confusing. Please rewrite it more concisely and clearly.

Response: Revised as the following (line 106 to 109)

"The poor understanding on mixing state of BC and BrC in individual particles will prevent the further simulation of atmospheric climate and aerosol-cloud interaction in the Arctic through the

current atmospheric models (Browse et al., 2013; Samset et al., 2014; Zanatta et al., 2018)."

9 Line 99-100: Change "collected on 7 to 23 August, 2012 in the Arctic." to "collected in the Artic between August 7th and 23rd, 2012." Response: Corrected.

10 Line 104: replace "on substrate" with "on a substrate" Response: deleted the word here

11 Line 118: Change "samples between 7 and 23 August, 2012." to "samples collected between August 7th and 23rd, 2012." Response: Corrected

12 Line 119: Replace "analyzed for TEM analysis" with "analyzed with TEM" Response: Revised

13 114-131: Restructure your sampling section. As written the reader may think that you sampled with 2 separate samplers an individual particle sampler and cascade impactor. After reading the paper, there is only one sampler. This confusion can be remedied by introducing the cascade impactor earlier.

Response: Revised the part as below (line 133 to 136)

"A sampler containing a single-stage impactor with a 0.5-mm-diameter jet nozzle (Genstar Electronic Technology, China) was used to collect individual particles by the air flow rate at 1.5 l min-1. Aerosol particles were collected onto copper TEM grids coated with carbon film."

14 122-123: Add the top size cutoff for this sampler. Response: Added

15 138: replace "within a" with "for a" Response: Corrected

16 148: 2002 particles examined over all the samples, or in a specific filter? Response: Added the Table S1 in the sentence. Table S1 can show how many particles we analyzed in each sample.

17 151: Clarify this sentence. Do you mean to say you ", we only checked elemental compositions of 20-30 particles" in each sample?

Response: We added more detailed description (line 165 to 180)

"In the clean Arctic air, there are simply particle types including sea salt, sulfate, soot, and OM. Because soot particles have chain-like aggregation, it is not necessary to check their elemental composition. Sea salt particles display spherical or square shapes and are stable under the electron beam in TEM but sulfate particles are spherical but flats on the substrate and produce unstable bubble under the electron beam (Buseck and Posfai, 1999; Chi et al., 2015). TEM observations also can clearly identify sulfate particles or sulfate with OM coating. Therefore, we

can easily identify Arctic particle types based on their morphology. Because of the time-consuming in the experiment, it is not necessary to frequently check elemental composition of the same particle type. For the data statistic in this study, we randomly checked elemental composition of 20-30 particles in each sample (Table S1). EDS spectra of 575 particles were manually selected and saved in the computer for elemental composition analysis. Particles examined by TEM were dry at the time of observation in the vacuum of the electron microscope. In our study, the effects of water and other semi-volatile organics were not considered as they evaporate in the vacuum."

18 155-156: The statement about Cu is redundant. This has already been stated in line 120 and 140.

Response: Corrected:

19 156-157: What is the difference between what is stated here and what is stated in lines 147-149?

Response: In lines 147-149 descript how to analyze particles on the substrate using TEM/EDS. In 156-157, we used scanning TEM. The method is one other function in the TEM which can give element profile along with the one line or elemental mapping in the targeted individual particle. We added more description here and tell the reader why we did the STEM here.

The following paragraph was revised to (line 181 to 188)

" Elemental mapping and line profile of individual aerosol particles were obtained from the EDX scanning operation mode of TEM (STEM). The STEM information can clearly display elemental distribution in the targeted individual particles which cannot be provided by the above EDS examination. Based on preliminary individual analysis, we further chose the typical samples containing abundant sulfate with OM coating for the STEM analysis. The high-resolution details of elemental distribution in individual particles can further prove the details of the mixing structure of sulfate and OM in individual particles."

20 162-163: Replace "is the image analysis platform..." with "is an image analysis platform". Response: Corrected

21 177: Replace "Organic Matters" with "Organic Matter" Response: Corrected

22 Line 184: Replace "TEM grids was" with either "The TEM grid was" or "TEM grids were" Response: Corrected

23 Line 192-194: Incomplete sentence Response: Corrected "AFM with a digital nanoscope IIIa instrument operating in the tapping mode was used to observe surface morphology of individual aerosol particles and measure particle thickness."

24 Line 194: I stopped making basic grammar and structure critiques at this point.

Response: Thank you very much. We carefully checked them and improved the English writing. Please see the red markers in the revised ms.

25 Line 222: Does treating this as a core-shell system with BC in the middle and sulfate and OM mixed on the outside have any basis? You've indicated that you have soot inclusions on the outside of predominantly sulfate particles, so why would soot be on the inside? Response: We've deleted the part with soot – see response in comment above.

26 Line 224: This sentence reads as if you've calculated the refractive index of the particles. Did you measure the optical properties of these particles? Response: We didn't measure their optical properties. We deleted this sentence and revised the part.

27 Line228: "In this study". It is unclear whether this is referring to your previous work or this manuscript.

Response: corrected

28 Line 253: Is it possible that coagulation of primary organic particles and S-rich particles could have led to the formation of organic coatings? Are you sure assumption that all organic coatings are secondary valid?

Response: The possibility could happen in the polluted air due to high particle number. As our previous study, Chi et al., (2017) determined sea salt particles collected in the Arctic air. We didn't observe the association of sulfate and sea salts particles, although we found many aged sea salt particles. The observations again suggest the particle coagulation is unlikely to be important in the Arctic air.

29 Line 272: Can you say a percentage of NSS particles that are S-rich with an OM coating? Or a percentage of S-rich particles that have a coating? This would strengthen the paper if an actual number was given.

Response: No, we cannot say NSS particles are S-rich with OM coating particles. Figure S4 shows 39% by number of all the analyzed particles were NSS-particles but 29% particles contain sulfate. Here we found 73% of the analyzed NSS-particles are S-rich with OM coating. We've changed this

line 310 to 312: "Here we focused on S-rich, soot, and OM particles as the major non-sea salt particle (NSS-particle, 39±5%) in the analyzed samples, which are approximately 29±7% of 2002 particles (Figure 3). "

L323-324: "A majority of 781 analyzed NSS-particles (74% by particle number) have a sulfate core and OM coating (Figures 4 and 5)."

30 Line 274-276: Are these specific samples special or is there something that you think may have caused the low frequency of soot inclusions? If so, please explain why. Response: We only found some fresh soot particles in three samples out of the 21 samples. We found many sulfate with soot inclusions. Previous works showed that BC is very low in the Arctic. We've changed this (line 326 to 333).

"The mixing structure is different from our previous findings in polluted air that soot is normally mixed with sulfate instead of OM coating (Li et al., 2016). Moreover, we noticed that a few chain-like soot aggregates (1.3% in all analyzed particles) (Figure S5) only occurred in three samples during the sampling period (Table S1). Considering the remoteness of the sampling site, such fresh soot particles are likely to be of local origin, including shipping and flaring (Gilgen et al., 2018; Peters et al., 2011). Indeed, we found a few of ships moving in Arctic Ocean during these days from the Ny-Ålesund town."

31 Line 277: This statement needs to be better supported. Just because a site is remote does not mean particles are local in origin. If this is supported by your trajectory calculations, mention them here. Also, don't the soot inclusions also imply that perhaps the OM is not secondary in nature? Soot is 100% primary and often co-emitted with primary OM, so if there is soot associated with OM coatings the soot itself is primary and so possibly some of the OM is primary as well.

Response: We found fresh soot particles in Figure 6. Indeed, TEM image shows that the soot particles have very thin OM coating.

It is true that soot is often co-emitted with primary OM. However, there are plenty of evidence to show that the primary soot aggregate particles are not heavily mixed with OM and in particularly not coated with thick OM (Wang et al., 2017). Furthermore, It is well known that SOA is dominant in the remote air (Jimenez et al., 2009).

32 Line 280: You need to define what satellite particles are; you have not discussed or defined them previously. Are they simply splatter of liquid portions of the particle when the particle is collected?

Response: Thanks. We define it.

L334-337 "TEM observations showed that some sulfate particles had unique morphology that a sulfate particle was surrounded by some smaller particles (Figure 5a). They are often called "satellite" particles as they were distributed from the central particles when impacted on the substrate during sample collection."

33 Line 280-281: is there a reason why satellite particles would have been observed on these days but not other days?

Response: The samples collected during 9-15, August contained abundant sulfate particles, OM, and soot particles (Table S1). Also, Figure 2 shows that air masses from North American during these days were mainly dominant. The reasons should depend on their air masses. Although the mechanism of this point is interesting, we could not give specific evidence under the current data.

34 Line 281: This is misleading and implies you performed the NanoSIMS analysis on 11 samples. In the methods section, it says only two samples were analyzed with NanoSIMS. Response: Corrected

L339-342: "NanoSIMS analysis further provided more information that the satellite particles

selected from the samples (Table S1) have strong 32S- (Figure 6a, c) and 16O- signals (Figure 6d) as well as weak 12C14N- signals (Figure 6a, b)."

35 Line 287-289: This is misleading. It reads as if you have done molecular characterization of the organic matter.

Response: Corrected

L346-347 "Indeed, Fu et al. (2008) found that polyacids are the most abundant organic compounds, followed by phthalates, aromatic acids, and fatty acids in Arctic aerosol particles."

36 Line 314: You back trajectories are only for specific days, be transparent about this in the Discussion

Response: Corrected. Add one Figure S2

37 Line 317-318: This conflicts with your earlier comment that most BC should be local. Response: The result does not conflict with earlier comments. We only pointed out the fresh soot particles (only 1.3%) may come from local shipping or combustion activitgies. It doesn't mean most BC were from local sources. In this study, we did not determine where the BC comes from.

38 Line 326-342: This discussion reads like a list of facts, but why they are all relevant is not always stated. Explain why each observation is important and how it adds context to your results. Response: We carefully revised the part as below

L338-408: "The sulfate core-OM shell structure observed in the Arctic summer atmosphere is similar to those in the background or rural air in other places (Li et al., 2016; Moffet et al., 2013). Based on the images from electron microscopies, we can infer that OM coating thickness in the arctic atmosphere was comparable with them in rural places but higher than them in urban places. During the transports, organic coatings on sulfates were considered as the secondary organic aerosols and their masses increase following particle aging and growth (Li et al., 2016; Moffet et al., 2013; Sierau et al., 2014). Figures 1 and 2 show that most of particles in the air masses transported long distance from North American. The result indicates that these long-range transportation of secondary sulfate particles have enough time to experience the possible atmospheric heterogeneous reactions on particle surfaces or cloud processes in the Arctic air. Similarly, Moffet et al. (2013) found that soot inclusions occurred in OM coating when OM coating on sulfates built up through photochemical activity and pollution buildup the Sacramento urban plume aged. On the other hand, the sulfate/OM particles with soot inclusions are probably formed in a similar way as those found elsewhere (Li et al., 2016) - e.g., soot particles may have acted as nuclei for secondary sulfate or organic uptake during their transports (Riemer et al., 2009). Similarly, besides the OM coating in the Arctic particles, Leck and Svensson (2015) found some biogenic aerosols like gel-aggregate containing bacterium in ultrafine particles. However, we didn't find any gel-like particles in the samples because our sampler had very low efficiency for ultrafine particles."

39 Line 343: Why is dry included here? Are there also wet particles that you have not discussed?

Response: Deleted the "Dry" here

40 Line 344-347: Explain LLPS in simpler terms and why it's important. Response: Revised the part.

L409-428: "TEM images show that most of the internally mixed sulfate particles display sulfate core and OM coating on the substrate (Figures 4a and 5b, c). The sulfate and OM separation in individual particles were defined by You et al. (2012) as liquid-liquid phase separation (LLPS). Concerning the knowledges of the LLPS can better understand particle hygroscopicity, heterogeneous reactions between reactive gases on particle surface, and organic aging (You et al., 2012). They also reported that the LLPS can reflect the O:C ratio in the OM, which is roughly \leq 0.5. In this study, we did observe the LLPS in almost all the fine sulfate particles, which indicates that the secondary OM in the coating might be not highly aged. Therefore, we speculate that the thick OM coatings were consistently built up during the long range transport of sulfate particles and part of secondary OM in the coating likely formed in Arctic area. Indeed, some studies reported that there are various sources of organic precursors during the Arctic area, such as biogenic VOCs from ice melting and open water (Dall'Osto et al., 2017) and anthropogenic VOCs from shipping emissions in summertime (Gilgen et al., 2018). The dependence of OM volume on particle size (Figure S6) suggests that the suspended sulfate particles are initially important surface for secondary OM formation. Moreover, the common OM coating on sulfate particles indicates that secondary OM as the surfaces of fine particles might govern the possible heterogeneous reactions between reactive gases and sulfate particles in the Arctic air."

41 Line 346: It's unclear what 90% is referring to in this statement. Response: We deleted this sentence here

42 Line 348: There's likely a complex relationship between phase state, oxidation state, and humidity. This needs to be investigated and explained further if statements about aerosol age are going to be made. Additionally, shouldn't you see a variety of ages of aerosol in your samples? Showing contrast between aged and unaged particles would be interesting and convincing.

Response: We re-wrote this part. We pointed out that the implications including LLSP in the Arctic air. Almost all particles we observed the aged. We do not have a parameter to define the age of the aerosol population and we are unable to answer the question clearly.

43 Line 376-378: This is circular reasoning because 12C14N- was what you used to identify OM so of course it was observed in the OM coatings. I don't think this data set is appropriate to make conclusions about the N content of OM coatings. That said, if you were able to calculate the mass concentration of N in the coatings with NanoSIMS that might give you a better indication of the BrC potential of the OM.

Response: It is impossible to calculate the N concentration in the coatings based on the NanoSIMS data here. The literature as the below proved that CN- can indicate N-containing organic matter.

Herrmann, A. M., K. Ritz, N. Nunan, P. L. Clode, J. Pett-Ridge, M. R. Kilburn, D. V. Murphy, A. G.

O'Donnell, and E. A. Stockdale (2007), Nano-scale secondary ion mass spectrometry — A new analytical tool in biogeochemistry and soil ecology: A review article, Soil Biology and Biochemistry, 39(8), 1835-1850.

44 Line 390-393: The average absorption cross-section is reported on a particle basis. This would be much more meaningful if it was extrapolated to environmental conditions. Because you've sampled from the atmosphere, you should be able to approximate particle concentrations, correct? Is so, you could back calculate this to an actual atmospheric absorption contribution and compare it to expected absorption from other species and measurements. This would be significantly more meaningful.

Response: Yes, we agree it would be useful to extrapolate to environmental conditions but as stated by the reviewers, we do not want to over-interpret our results. We cannot extrapolate because we do not have ambient particle concentration data. One single paper will not be able to answer all questions and we do suggest for future work to answer these questions.

45 Line 413, 260: 29% number should include a standard deviation. Response: added

46 Line 410: The last section shouldn't simply repeat what was stated in the above sections, but instead present the data in additional context and discuss the implications. Response: We made major revision in the section. Please see the revised red words. References:

- Abbatt, J.P.D., Leaitch, W.R., Aliabadi, A.A., Bertram, A.K., Blanchet, J.P., Boivin-Rioux, A., Bozem, H., Burkart, J., Chang, R.Y.W., Charette, J., Chaubey, J.P., Christensen, R.J., Cirisan, A., Collins, D.B., Croft, B., Dionne, J., Evans, G.J., Fletcher, C.G., Galí, M., Ghahremaninezhad, R., Girard, E., Gong, W., Gosselin, M., Gourdal, M., Hanna, S.J., Hayashida, H., Herber, A.B., Hesaraki, S., Hoor, P., Huang, L., Hussherr, R., Irish, V.E., Keita, S.A., Kodros, J.K., Köllner, F., Kolonjari, F., Kunkel, D., Ladino, L.A., Law, K., Levasseur, M., Libois, Q., Liggio, J., Lizotte, M., Macdonald, K.M., Mahmood, R., Martin, R.V., Mason, R.H., Miller, L.A., Moravek, A., Mortenson, E., Mungall, E.L., Murphy, J.G., Namazi, M., Norman, A.L., O'Neill, N.T., Pierce, J.R., Russell, L.M., Schneider, J., Schulz, H., Sharma, S., Si, M., Staebler, R.M., Steiner, N.S., Thomas, J.L., von Salzen, K., Wentzell, J.J.B., Willis, M.D., Wentworth, G.R., Xu, J.W., Yakobi-Hancock, J.D.: Overview paper: New insights into aerosol and climate in the Arctic, Atmos. Chem. Phys., 19 (4), 2527-2560, 2019.
- Bohren, C.F., Huffman, D.R., 1983. Absorption and scattering of light by small particles. John Wiley & Sons, Inc., New York, USA.
- Browse, J., Carslaw, K.S., Schmidt, A., Corbett, J.J.: Impact of future Arctic shipping on high-latitude black carbon deposition, Geophys. Res. Lett.,, 40 (16), 4459-4463, 2013.
- Buseck, P.R., Posfai, M.: Airborne minerals and related aerosol particles: Effects on climate and the environment, P. Natl. Acad. Sci. USA, 96 (7), 3372-3379, 1999.
- Chi, J.W., Li, W.J., Zhang, D.Z., Zhang, J.C., Lin, Y.T., Shen, X.J., Sun, J.Y., Chen, J.M., Zhang, X.Y., Zhang, Y.M., Wang, W.X.: Sea salt aerosols as a reactive surface for inorganic and organic acidic gases in the Arctic troposphere, Atmos. Chem. Phys., 15 (19), 11341-11353, 2015.
- Dall'Osto, M., Beddows, D.C.S., Tunved, P., Krejci, R., Ström, J., Hansson, H.C., Yoon, Y.J., Park, K.-T., Becagli, S., Udisti, R., Onasch, T., O'Dowd, C.D., Simó, R., Harrison, R.M.: Arctic sea ice melt leads to atmospheric new particle formation, Sci. Rep., 7 (1), 3318, 2017.
- Fu, P., Kawamura, K., Barrie, L.A.: Photochemical and Other Sources of Organic Compounds in the Canadian High Arctic Aerosol Pollution during Winter–Spring, Environ. Sci. Technol., 43 (2), 286-292, 2008.
- Gilgen, A., Huang, W.T.K., Ickes, L., Neubauer, D., Lohmann, U.: How important are future marine and shipping aerosol emissions in a warming Arctic summer and autumn?, Atmos. Chem. Phys., 18 (14), 10521-10555, 2018.
- Jimenez, J.L., Canagaratna, M.R., Donahue, N.M., Prevot, A.S.H., Zhang, Q., Kroll, J.H., DeCarlo, P.F.,
 Allan, J.D., Coe, H., Ng, N.L., Aiken, A.C., Docherty, K.S., Ulbrich, I.M., Grieshop, A.P., Robinson,
 A.L., Duplissy, J., Smith, J.D., Wilson, K.R., Lanz, V.A., Hueglin, C., Sun, Y.L., Tian, J., Laaksonen,
 A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J.M.,
 Collins, D.R., Cubison, M.J., E., Dunlea, J., Huffman, J.A., Onasch, T.B., Alfarra, M.R., Williams,
 P.I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K.,
 Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun,
 J.Y., Zhang, Y.M., Dzepina, K., Kimmel, J.R., Sueper, D., Jayne, J.T., Herndon, S.C., Trimborn,
 A.M., Williams, L.R., Wood, E.C., Middlebrook, A.M., Kolb, C.E., Baltensperger, U., Worsnop,
 D.R.: Evolution of Organic Aerosols in the Atmosphere, Science, 326 (5959), 1525-1529, 2009.
- Leck, C., Svensson, E.: Importance of aerosol composition and mixing state for cloud droplet activation over the Arctic pack ice in summer, Atmos. Chem. Phys., 15 (5), 2545-2568, 2015.

- Li, W., Sun, J., Xu, L., Shi, Z., Riemer, N., Sun, Y., Fu, P., Zhang, J., Lin, Y., Wang, X., Shao, L., Chen, J., Zhang, X., Wang, Z., Wang, W.: A conceptual framework for mixing structures in individual aerosol particles, Journal of Geophysical Research: Atmospheres, 121 (22), 13,784-713,798, 2016.
- Moffet, R.C., Rödel, T.C., Kelly, S.T., Yu, X.Y., Carroll, G.T., Fast, J., Zaveri, R.A., Laskin, A., Gilles, M.K.: Spectro-microscopic measurements of carbonaceous aerosol aging in Central California, Atmos. Chem. Phys., 13 (20), 10445-10459, 2013.
- Peters, G.P., Nilssen, T.B., Lindholt, L., Eide, M.S., Glomsrød, S., Eide, L.I., Fuglestvedt, J.S.: Future emissions from shipping and petroleum activities in the Arctic, Atmos. Chem. Phys., 11 (11), 5305-5320, 2011.
- Riemer, N., West, M., Zaveri, R.A., Easter, R.C.: Simulating the evolution of soot mixing state with a particle resolved aerosol model, J. Geophys. Res., 114, doi:10.1029/2008JD011073, 2009.
- Samset, B.H., Myhre, G., Herber, A., Kondo, Y., Li, S.M., Moteki, N., Koike, M., Oshima, N., Schwarz, J.P., Balkanski, Y., Bauer, S.E., Bellouin, N., Berntsen, T.K., Bian, H., Chin, M., Diehl, T., Easter, R.C., Ghan, S.J., Iversen, T., Kirkevåg, A., Lamarque, J.F., Lin, G., Liu, X., Penner, J.E., Schulz, M., Seland, Ø., Skeie, R.B., Stier, P., Takemura, T., Tsigaridis, K., Zhang, K.: Modelled black carbon radiative forcing and atmospheric lifetime in AeroCom Phase II constrained by aircraft observations, Atmos. Chem. Phys., 14 (22), 12465-12477, 2014.
- Seinfeld, J., Pandis, S., 2006. Atmospheric Chemistry and Physics: From air pollution to climate change (2nd ed.). 1-1203 pp., John Wiley & Son, Inc., Hoboken, New Jersey.
- Sierau, B., Chang, R.Y.W., Leck, C., Paatero, J., Lohmann, U.: Single-particle characterization of the high-Arctic summertime aerosol, Atmos. Chem. Phys., 14 (14), 7409-7430, 2014.
- Wang, Y., Liu, F., He, C., Bi, L., Cheng, T., Wang, Z., Zhang, H., Zhang, X., Shi, Z., Li, W.: Fractal Dimensions and Mixing Structures of Soot Particles during Atmospheric Processing, Environ. Sci. Tech. Let., 4 (11), 487-493, 2017.
- You, Y., Renbaum-Wolff, L., Carreras-Sospedra, M., Hanna, S.J., Hiranuma, N., Kamal, S., Smith, M.L., Zhang, X., Weber, R.J., Shilling, J.E., Dabdub, D., Martin, S.T., Bertram, A.K.: Images reveal that atmospheric particles can undergo liquid–liquid phase separations, Proceedings of the National Academy of Sciences of the United States of America 109 (33), 13188-13193, 2012.
- Zanatta, M., Laj, P., Gysel, M., Baltensperger, U., Vratolis, S., Eleftheriadis, K., Kondo, Y., Dubuisson, P., Winiarek, V., Kazadzis, S., Tunved, P., Jacobi, H.W.: Effects of mixing state on optical and radiative properties of black carbon in the European Arctic, Atmos. Chem. Phys., 18 (19), 14037-14057, 2018.
- Jiang, H., A. L. Frie, A. Lavi, J. Y. Chen, H. Zhang, R. Bahreini, and Y.-H. Lin (2019), Brown Carbon Formation from Nighttime Chemistry of Unsaturated Heterocyclic Volatile Organic Compounds, Environmental Science & Technology Letters, DOI: 10.1021/acs.estlett.1029b00017.
- Yang, Y., Wang, H., Smith, S.J., Easter, R.C., Rasch, P.J.: Sulfate Aerosol in the Arctic: Source Attribution and Radiative Forcing, Journal of Geophysical Research: Atmospheres, 123 (3), 1899-1918, 2018.

1	Organic coating on sulfate and soot particles in summer Arctic
2	atmosphere
3	
4	Hua Yu ^{1,2} , Weijun Li ^{2*} , Yangmei Zhang ³ , Peter Tunved ⁴ , Manuel Dall'Osto ⁵ , Xiaojing
5	Shen ³ , Junying Sun ³ , Xiaoye Zhang ³ , Jianchao Zhang ⁶ , Zongbo Shi ^{7,8*}
6	
7	¹ College of Life and Environmental Sciences, Hangzhou Normal University, 310036, Hangzhou, China
8	² Department of Atmospheric Sciences, School of Earth Sciences, Zhejiang University, 310027,
9	Hangzhou, China
10	³ Key Laboratory of Atmospheric Chemistry, Chinese Academy of Meteorological Sciences, Beijing,
11	China
12	⁴ Department of Environmental Science and Analytical Chemistry, Stockholm University, 10691,
13	Stockholm, Sweden
14	⁵ Institute of Marine Sciences, ICM-CSIC, Passeig Mar fim de la Barceloneta, 37-49, E-08003,
15	Barcelona, Spain
16	⁶ Key Laboratory of the Earth's Deep Interior, Institute of Geology and Geophysics, Chinese Academy
17	of Sciences, 100029, China
18	⁷ School of Geography, Earth and Environmental Sciences, University of Birmingham, Birmingham,
19	UK
20	⁸ Institute of Surface Earth System Science, Tianjin University, Tianjin, China
21	
22	[*] Corresponding Emails: liweijun@zju.edu.cn; z.shi@bham.ac.uk
23	

24 Abstract

Interaction of anthropogenic particles with radiation and clouds plays an important 25 role on Arctic climate change. Mixing state of aerosols is a key parameter to influence 26 aerosol-cloud and aerosol-radiation interaction. However, little is known on this 27 parameter in the Arctic, preventing an accurate representation of this information in 28 29 global models. Here we used transmission electron microscopy with energy-dispersive X-ray spectrometry (TEM/EDS), scanning TEM, scanning electron 30 31 microscopy (SEM), nanoscale secondary ion mass spectrometry (NanoSIMS), and atomic forces microscopy (AFM) to determine the size and mixing properties of 32 individual particles at 100 nm - 10 μ m, with a particular focus on sulfate and 33 carbonaceous particles. We found that non-sea salt sulfate particles with size range at 34 100-2000 nm were commonly coated with organic matter (OM) in summer. 20% of 35 sulfate particles also had soot inclusions which only appeared in the OM coating. The 36 OM coating is estimated to contribute to 63% of the particle volume on average. To 37 understand how OM coating influences optical properties of sulfate particles, the Mie 38 39 theory of the core-shell model was applied to calculate optical properties of individual sulfate particles. The result shows that absorption cross section (ACS) of individual 40 OM-coated particles significantly increased when assuming the OM coating as 41 light-absorbing brown carbon (BrC) and the ACS also increased following the 42 increasing particle size. The microscopic observations suggest that OM modulates the 43 mixing structure of fine Arctic sulfate particles, which may determine their 44 hygroscopicity and optical properties. 45

46

47 **1. Introduction**

Surface temperatures are rising faster in the Arctic than the rest of globe (IPCC, 48 2013). Although increased human-induced emissions of long-lived greenhouse gases 49 are certainly one of the driving factors, air pollutants, such as aerosols and ozone, are 50 also important contributors to climate change in the Arctic (Law and Stohl, 2007; 51 52 Shindell, 2007). It is well known that aerosols from northern mid-altitude continents affect the sea ice albedo by altering the heat balance of the atmosphere and surface 53 54 (Hansen and Nazarenko, 2004; Jacob et al., 2010; Shindell, 2007). These aerosols in Arctic atmosphere include sea salt, sulfate, particulate organic matter (OM), and to a 55 lesser extent, ammonium, nitrate, black carbon (BC) (Hara et al., 2003; Quinn et al., 56 2007) and mineral dust particles (Dagsson-Waldhauserova et al., 2013). Studies show 57 BC in the Arctic absorbs solar radiation in the atmosphere and when deposited on 58 snow (Iziomon et al., 2006; Koch and Hansen, 2005; Sand et al., 2013; Shindell, 59 2007). Moreover, Maahn et al. (2017) used aircraft in situ observation of clouds and 60 aerosols and found that concentration of BC are enhanced below the clouds in the 61 62 Arctic and further influence the mean effective radii of cloud droplets which lead to the suppressed drizzle production and precipitation. 63

BC, commonly called "soot" is derived from the combustion sources such as 64 diesel engines, residential solid fuel, and open burning (Bond et al., 2013). Some 65 studies investigated the possible sources of these BC particles, including natural gas 66 flaring (Qi et al., 2017) and ship emissions in the Arctic (Browse et al., 2013; 67 Weinbruch et al., 2012) and emissions of biomass burning and fossil fuels in the 68 northern hemisphere (Winiger et al., 2016; Xu et al., 2017). For example, Winiger et 69 70 al.(2017) showed that most Arctic BC is sourced from domestic activities (35%) and transportation (38%), with only minor contributions from gas flaring (6%), power 71 72 plants (9%), and open fires (12%).

Accumulation of secondary organic aerosols, a significant fraction of the new particles grows to sizes that are active in cloud droplet formation in the Arctic (Abbatt et al., 2019). More than 100 organic species have been detected in the Arctic aerosols and polyacids are the most abundant compound class, followed by phthalates, 77 aromatic acids, fatty acids, fatty alcohols, sugars/sugar alcohols, and n-alkanes (Fu et al., 2008). Recently, certain organic aerosols, referred to as brown carbon (BrC), have 78 been recognized as an important light-absorbing carbonaceous aerosol after BC in the 79 troposphere (Alexander et al., 2008; Andreae and Gelencser, 2006; Feng et al., 2013; 80 Lack et al., 2012). BrC can be directly emitted from combustion sources or formed in 81 the atmosphere via photo-chemical aging (Jiang et al., 2019; Saleh et al., 2013; 82 Updyke et al., 2012). Moreover, aging of secondary organic aerosols can significantly 83 84 contribute to BrC during atmospheric transports (Laskin et al., 2015). Feng et al.(2013) estimated that on average, BrC accounts for 66% of total OM mass globally and its 85 light absorption is about 26% of BC. 86

BC and BrC are often internally mixed with other non-absorbing aerosols, such as 87 sulfate (Lack et al., 2012; Laskin et al., 2015). Internal mixing means that a single 88 particle simultaneously contains two or more types of aerosol components (Li et al., 89 2016). This internal mixing can enhance BC absorption by a factor of up to two (Bond 90 et al., 2013) and change the activity of cloud condensation nuclei (CCN) in the Arctic 91 92 atmosphere (Leck and Svensson, 2015; Martin et al., 2011). Spatial and temporal variations of aerosol composition, size distribution, and sources of Arctic aerosols 93 have been studied extensively in numerous ground-based, ship, airborne observations, 94 95 and various atmospheric models (Brock et al., 2011; Burkart et al., 2017; Chang et al., 2011; Dall Osto et al., 2017; Fu et al., 2008; Hara et al., 2003; Hegg et al., 2010; 96 Iziomon et al., 2006; Karl et al., 2013; Lathem et al., 2013; Leck and Bigg, 2008; 97 Leck and Svensson, 2015; Moore et al., 2011; Raatikainen et al., 2015; 98 Wöhrnschimmel et al., 2013; Winiger et al., 2017; Yang et al., 2018; Zangrando et al., 99 100 2013). A few previous studies also looked at the mixing states of coarse aerosol particles in Arctic troposphere (Behrenfeldt et al., 2008; Chi et al., 2015; Geng et al., 101 2010; Hara et al., 2003; Leck and Svensson, 2015; Moroni et al., 2017; Raatikainen et 102 al., 2015; Sierau et al., 2014), but those of fine non-sea salt particles, including the 103 most important short-lived climate forcers - BC and BrC (Feng et al., 2013; Fu et al., 104 105 2008; Kirpes et al., 2018; Laskin et al., 2015; Leck and Svensson, 2015), are poorly characterized. The poor understanding on mixing state of BC and BrC in individual 106

particles will prevent the further simulation of atmospheric climate and aerosol-cloud
interaction in the Arctic through the current atmospheric models (Browse et al., 2013;
Samset et al., 2014; Zanatta et al., 2018).

In this study, individual aerosol particles were collected in the Arctic during 7-23 August, 2012. We combined the data from various microscopic instruments to systematically determine the size, composition, and mixing properties of individual particles, with a particular focus on sulfate and carbonaceous particles. Mie theory was used to test how OM coating influences optical properties of sulfate particles in the Arctic when OM was assumed as BrC. The results are discussed in the context of aerosol-radiation and cloud interaction.

117

118 **2. Experimental section**

119 2.1 Field campaign

The Svalbard archipelago includes all landmasses between 74 and 81 degrees North and 10 and 35 degrees East (Figure 1). The islands cover 63000 km². Ny-Ålesund town is situated on the west coast of the largest island, Spitsbergen. Ny-Ålesund town is situated only 1200 km from the North Pole and represents a central platform for Arctic research. The sampling place represents remote Arctic conditions.

An individual particle sampler at Chinese Arctic Yellow River Station (78°55'N, 126 11°56'E) collected individual particles (Chi et al., 2015; Geng et al., 2010). The 127 sampling site is about 2 km far away from the Zeppelin observatory station (78.9N 128 by the Ny-Ålesund Science 129 11.88E) running Managers Committee 130 (https://www.esrl.noaa.gov/psd/iasoa/stations/nyalesund). Two to three samples were regularly collected at 9:00, 16:00, 21:00 (local time) of each day, with a total of 46 131 samples during 7-23 August, 2012. 132

A sampler containing a single-stage impactor with a 0.5-mm-diameter jet nozzle (Genstar Electronic Technology, China) was used to collect individual particles by the air flow rate at 1.5 l min⁻¹. Aerosol particles were collected onto copper TEM grids coated with carbon film. This sampler has a collection efficiency of 31% at 100 nm

aerodynamic diameter and 50% at 200 nm if the density of the particles is 2 g cm⁻³. 137 The sampler can collect particles with $< 10 \mu m$ aerodynamic diameter on TEM grids. 138 Sampling times varied from twenty minutes to two hours in clean remote Arctic area. 139 After collection, each sample was placed in a sealed dry plastic tube and stored in a 140 desiccator at 20 \pm 3% RH for analysis. Ambient laboratory conditions (17–23% RH 141 and 19–21 °C) is effective at preserving individual hygroscopic aerosol particles and 142 reducing changes that would alter samples and subsequent data interpretation 143 144 (Laskina et al., 2015). The sample information such as local sampling date and time and meteorological conditions (e.g., temperature (T), relative humidity (RH), pressure 145 (P), wind direction (WD), wind speed (WS)) are listed in Table S1. 146

147

148 **2.2 TEM measurement**

Individual particle samples were examined by a JEOL JEM-2100 transmission 149 electron microscopy operated at 200 kV with an energy-dispersive X-ray 150 spectrometry (TEM/EDS). TEM can observe the mixing structure of different aerosol 151 152 components within an individual particle on the substrate because electron beam transmit through the specimen to form an image. EDS spectra were acquired for a 153 maximum time of 30 s to minimize potential beam damage and collect particle X-ray 154 spectra with sufficient intensity. TEM grids are made of copper (Cu) and covered by a 155 carbon-reinforced substrate, so Cu is excluded from the quantitative analyses of the 156 particles. Because of the substrate contribution, C content in TEM grid coated by 157 carbon film might be overestimated in EDS spectra of individual particles. 158

The distribution of aerosol particles on TEM grids was not uniform, with coarser 159 160 particles occurring near the center and finer particles on the periphery. Therefore, to ensure that the analyzed particles are representative, five areas were chosen from the 161 center and periphery of the sampling spot on each grid. Through a labor-intensive 162 operation, 2002 aerosol particles with diameter $< 10 \mu m$ in 21 samples were analyzed 163 by TEM/EDS (Table S1). To check elemental composition of individual particles, 164 EDX was manually used to obtain EDS spectra of individual particles. In the clean 165 Arctic air, there are simply particle types including sea salt, sulfate, soot, and OM. 166

Because soot particles have chain-like aggregation, it is not necessary to check their 167 elemental composition. Sea salt particles display spherical or square shapes and are 168 stable under the electron beam in TEM but sulfate particles are spherical but flats on 169 the substrate and produce unstable bubble under the electron beam (Buseck and Posfai, 170 1999; Chi et al., 2015). TEM observations also can clearly identify sulfate particles or 171 sulfate with OM coating. Therefore, we can easily identify Arctic particle types based 172 on their morphology. Because of the time-consuming in the experiment, it is not 173 174 necessary to frequently check elemental composition of the same particle type. For the data statistic in this study, we randomly checked elemental composition of 20-30 175 particles in each sample (Table S1). EDS spectra of 575 particles were manually 176 selected and saved in the computer for elemental composition analysis. Particles 177 examined by TEM were dry at the time of observation in the vacuum of the electron 178 179 microscope. In our study, the effects of water and other semi-volatile organics were not considered as they evaporate in the vacuum. 180

Elemental mapping and line profile of individual aerosol particles were obtained 181 182 from the EDX scanning operation mode of TEM (STEM). The STEM information can clearly display elemental distribution in the targeted individual particles which cannot 183 be provided by the above EDS examination. Based on preliminary individual analysis, 184 we further chose the typical samples containing abundant sulfate with OM coating for 185 the STEM analysis. The high-resolution details of elemental distribution in individual 186 particles can further prove the details of the mixing structure of sulfate and OM in 187 individual particles. 188

The iTEM software (Olympus soft imaging solutions GmbH, Germany) is an 189 190 image analysis platform for electron microscopy. In this study, it was used to manually or automatically obtain area, perimeter, and equivalent circle diameter 191 (ECD) of individual particles through identifying boundary of every particle in TEM 192 images. In these analyzed samples, we found there were abundant fine sulfate 193 particles in 11 samples collected during 9-15 August, 2012. In other samples, there 194 195 were only a few sulfate particles and more sea salt particles. Based on the TEM observations, we selected the samples containing more sulfate particles to further do 196

197 198 other microscopic analyses as below.

199 2.3 NanoSIMS measurement

Because the sulfate particles collected in the Arctic had good consistent property 200 (e.g., elemental composition and mixing state) from TEM observations, we just 201 selected three samples containing abundant fine sulfate particles (Table S1) for further 202 studies. These three samples listed in Table S1 were analyzed using a nanoscale 203 204 secondary ion mass spectrometry (NanoSIMS) 50L (CAMECA Instruments, Geneviers, France) instrument. A micro-cesium source was used to generate Cs⁺ 205 primary ions, with an impact energy of 16 kV for sample interrogation. The primary 206 beam was stepped across the sample to produce element specific, quantitative digital 207 images. The Cs⁺ primary ion beam was used to obtain ¹⁶O⁻, ¹²C¹⁴N⁻, ¹⁴N¹⁶O⁻, ³²S⁻, 208 35 Cl⁻, and 16 O²³Na⁻ ions in this study. The NanoSIMS analysis can obtain ion mapping 209 of particles with nanometer spatial resolution over a broad range of particle sizes 210 (Figure S1). Because the substrate of TEM grid is carbon, CN⁻ is adopted to represent 211 212 OM in individual particles (Chi et al., 2015; Ghosal et al., 2014). S⁻ is used to infer the presence of sulfates in individual particles (Li et al., 2017). Finally, the NanoSIMS 213 obtained ion mapping of 32 sulfate particles. 214

215

216 **2.4 SEM and AFM measurement**

Because TEM could not vertically observe OM coating and sulfate core, we 217 conducted one special experiment using a Zeiss ultra 55 scanning electron microscopy 218 (SEM) with EDS. The TEM grids were mounted onto an aluminum SEM stub and 219 directly observed in secondary electron image mode. SEM analysis was operated at 220 10 kv of extra high tension (EHT) and 9.7 mm of work distance (WD). Processes such 221 as sample moving, analysis region selection and imaging were controlled by computer. 222 The specimen stage in SEM was tilted at the range of $0-75^{\circ}$, and then we vertically 223 observed thickness of OM coating and sulfate core on the substrate. To verify vertical 224 225 property of individual S-rich particles impacting on the substrate, we observed two typical samples containing abundant sulfate particles using the SEM (Table S1). 226

227 AFM with a digital nanoscope IIIa instrument operating in the tapping mode was used to observe surface morphology of individual aerosol particles and measure 228 particle thickness. The tapping AFM has a cantilever and conical tip of 10 nm radius. 229 By using AFM, a general image of the particles is taken at 10 µm full scan size, which 230 generally includes 1-2 particles depending on the exact location. In this study, we are 231 only interested in the sulfate-containing particles. AFM provides surface information 232 and morphology of 17 particles but no composition. Samples were firstly quickly 233 234 examined by the TEM under low magnification mode. In case, the operation roughly identified S-containing particles and didn't damage the secondary sulfate particles 235 under the electron. Because TEM grids have coordinates letters, we can exactly find 236 the same particles on the substrate in AFM examined in TEM experiments. The 237 procedures can exclude sea salt particles in the AFM image. As a result, the same 238 samples observed by TEM were then examined in AFM to obtain 3-D image of 239 secondary sulfate particles and their volume. Because individual particles collected in 240 Arctic air were scattered on the substrate, we only obtained 17 effective data. After we 241 242 obtained AFM images of sulfate particles, the NanoScope analysis software can automatically obtain bearing area (A) and bearing volume (V) of each analyzed 243 particle according to the following formula. 244

245
$$A = \frac{4}{3}\pi r^2 = \frac{\pi d^2}{3} \to d = \sqrt{\frac{3A}{\pi}}$$
 (1)

246
$$V = \frac{4}{3}\pi r^3 = \frac{4}{3} \times \frac{\pi D^3}{8} \to D = \sqrt[3]{\frac{6V}{\pi}}$$
 (2)

Where x is the equivalent circle diameter (ECD) and y is the equivalent sphericaldiameter (ESD).

ECD of individual aerosol particles measured from the iTEM software can be further converted into ESD. Based on these data, we estimate one good linear correlation (y=0.38x) between ESD and ECD of sulfate particles impacting on the substrate. The value was further used to correct all the analyzed particles in TEM images (Chi et al., 2015).

254

255 **2.5 Calculation of BrC optical properties**

The refractive index used for the non-light-absorbing sulfate component was set to m=1.55 at 550 nm (Seinfeld and Pandis, 2006). The refractive index of OM (as BrC) is not known so we considered three scenarios: strongly absorbing (1.65-0.03i at 550 nm), moderately absorbing (1.65-0.003i at 550 nm), and non-absorbing OM (1.65 at 550 nm) (Feng et al., 2013). Although the refractive index has dependence on the wavelength between 350-870 nm, we tried to select the 550 nm as a case to test how OM coating influence sulfate particles in Arctic air.

BHCOAT Mie code by Bohren and Huffman (1983) was used to calculate the 263 optical properties, including scattering cross section (SCS), absorption cross section 264 (ACS), and single scattering albedo (SSA), assuming a core-shell structure. We firstly 265 calculated these parameters assuming a sulfate core and OM shell structure only 266 267 (ignoring some of the particles that contain soot core). Because the Mie code only can calculate the core-shell structure or homogeneous models, we assume sulfate as a core 268 and OM as a shell in individual particle to build the core-shell model. Based on the 269 270 core-shell standard mode (Li et al., 2016), we can calculate optical properties of individual internally mixed particles. 271

272 **2.6 Back trajectories of air masses and Lagrangian particle dispersion model**

Three-day (72 h) back trajectories of air masses were generated using a Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT) model at the Chinese Arctic Yellow River Station during August 2012, at an altitude of 500m above sea level (Figure 1). Most air masses originate in the Arctic Ocean, and are restricted to this vast marine region during the sampling periods. Based on the TEM observations, air masses from North America and Greenland brought abundant sulfate particles into the sampling area in summertime.

In order to determine the particle origins, the lagrangian particle dispersion model FLEXPART-WRF 3.1 (Brioude et al., 2013) was used. The FLEXPART-WRF model is using meteorological parameters from WRF dynamical simulation. The domain resolution is 50×50 km with 64 vertical levels. The FLEXPART-WRF simulations were launched in a backward mode over 10 days, with the Chinese Arctic Yellow

River Station as an origin. For each simulation (one per sample), 20000 285 pseudo-particles were released in a small volume around the station position. Each 286 single particle position evolution backward in time was determined by Lagrangian 287 dispersion calculation. Based on the TEM experiments and back trajectory of air 288 masses (Figure 1), we found that there were more S-rich with OM coating particles in 289 the samples collected on August 11, 12, 14 and 15, 2012. Therefore, we further did the 290 FLEXPART-WRF simulation of these four days (Figure 2). The emission intensity in 291 292 the Arctic area has been also shown in Figure S2.

293 **3. Results**

3.1 Composition and sources of aerosol particles

We summarized average elemental weight and frequency of individual Arctic 295 particles derived from the TEM/EDX. The result shows that O, Na, S, and Cl in 296 individual particles are dominant elements (Figure S3). On basis of the composition 297 and morphology of individual particles, we classified the particles into four major 298 groups: Na-rich (i.e. NaCl, Na₂SO₄, and NaNO₃), S-rich (i.e. ammonium sulfate and 299 300 sulfuric acid), and carbonaceous (soot and OM). The classification criteria of different particle types and their sources have been described in a separate study (Li et al., 301 2016). S-rich particles representing secondary inorganic particles (e.g., SO₄²⁻, NO₃⁻, 302 and NH_4^+) are transformed from gaseous SO₂, NO_x, and NH₃. OM can be divided into 303 primary organic matter (POM) and secondary organic matter (SOM). SOM is 304 produced from the chemical oxidation of volatile organic compounds (VOCs) and 305 often exhibits OM coating on S-rich particles. Na-rich particles in the marine air are 306 from sea spray and have typical near cubic shape. Soot particles, which contain C 307 308 with minor O, appear as a chain-like aggregate of carbon-bearing spheres. Our previous study well characterized aging mechanism of sea salt particles in summer 309 Arctic air (Chi et al., 2015). Here we focused on S-rich, soot, and OM particles as the 310 major non-sea salt particle (NSS-particle, $39\pm5\%$) in the analyzed samples, which are 311 approximately 29±7% of 2002 particles (Figure 3). 312

313

314 **3.2 OM coating on sulfate particles**

315 TEM observations revealed a common core-shell mixing structure in fine sulfate particles (Figure 4a). Elemental mapping of such internally mixed sulfate particles 316 shows C signals in the coating (C map, Figure 4b) and S and O signals in the center (S 317 and O map, Figure 4c, d). The elemental line profile of a sulfate particle also shows 318 sulfate core and C coating (Figure S4). Furthermore, ion maps of individual particles 319 from the NanoSIMS further exhibite ${}^{12}C^{14}N^{-}$ signals in the coating (red color in Figure 320 4e, f) and ${}^{32}S^{-}$ signals in the core (green color in Figure 4e, g). These results provide 321 strong evidence that the coating is OM and the core is sulfate. 322

A majority of 781 analyzed NSS-particles (74% by particle number) have a sulfate 323 core and OM coating (Figures 4 and 5). ~20% of them also contain small soot 324 inclusions but they only appeared in organic coating, rather than as the core mixed in 325 sulfate (Figure 5b). The mixing structure is different from our previous findings in 326 polluted air that soot is normally mixed with sulfate instead of OM coating (Li et al., 327 2016). Moreover, we noticed that a few chain-like soot aggregates (1.3% in all 328 analyzed particles) (Figure S5) only occurred in three samples during the sampling 329 330 period (Table S1). Considering the remoteness of the sampling site, such fresh soot particles are likely to be of local origin, including shipping and flaring (Gilgen et al., 331 2018; Peters et al., 2011). Indeed, we found a few of ships moving in Arctic Ocean 332 during these days from the Ny-Ålesund town. 333

TEM observations showed that some sulfate particles had unique morphology that 334 a sulfate particle was surrounded by some smaller particles (Figure 5a). They are 335 often called "satellite" particles as they were distributed from the central particles 336 when impacted on the substrate during sample collection. 16% of the analyzed sulfate 337 particles with satellite particles as shown in Figure 5a were detected in the samples 338 (Table S1) collected during 9-15 August. NanoSIMS analysis further provided more 339 information that the satellite particles selected from the samples (Table S1) have 340 strong ${}^{32}S^{-}$ (Figure 6a, c) and ${}^{16}O^{-}$ signals (Figure 6d) as well as weak ${}^{12}C^{14}N^{-}$ signals 341 (Figure 6a, b). The CN⁻ signal normally can represent organic aerosols (Chi et al., 342 343 2015; Ghosal et al., 2014). Previous studies showed that the similar satellite particles are normally considered as acidic sulfate (Buseck and Posfai, 1999; Iwasaka et al., 344

1983). Therefore, we can conclude that these acidic satellites not only contain sulfuric 345 acid but also some OM or organic acids. Indeed, Fu et al. (2008) found that polyacids 346 are the most abundant organic compounds, followed by phthalates, aromatic acids, 347 and fatty acids in Arctic aerosol particles. As a result, these Arctic sulfate particles 348 with satellites contain certain amounts of sulfuric or organic acids with liquid phase. 349 Back trajectories of air masses and FLEXPART both shows abundant sulfate particles 350 and some containing satellite particles were transported from Greenland and North 351 352 American (Figures 1 and 2).

AFM was used to obtain 3D image of individual secondary particles impacting on 353 the substrate. Figure 7a shows that the secondary particles normally have smooth 354 surface which is different from uneven surface of the Arctic fresh and aged NaCl 355 particles (Chi et al., 2015). Furthermore, we observed particle thickness through 356 tilting the specimen stage up to 75° in SEM. Figure 7a-b both shows that the 357 secondary particles look like thin pancake sticking on the substrate. Furthermore, the 358 sections of two secondary particles in the AFM images shows that the highest heights 359 360 of particles are only 0.15 (green line) and 0.26 (red line) of the corresponding horizontal size (Figure 7a). Here we can conclude that shape of individual particles 361 was modified when they impacted on the substrate following the airflow. Therefore, 362 the measured ECDs of individual particles in TEM images are much larger than the 363 real particle diameter. To calibrate the particle diameter, we obtained volume of dry 364 particles on the substrate and then calculated their equivalent sphere diameter (ESD) 365 in the AFM images (Figure 7c). ESD distribution of the secondary Arctic particles 366 displayed a peak at 340 nm, ranging from 100 nm to 2000 nm (Figure 7d). The core 367 368 particles, as sulfate or soot, had a peak at 240 nm and 120 nm, respectively (Figure 7d). In the core-shell particles, we knew size in all the analyzed particles and further 369 calculated volume of sulfate, OM, and/or soot within individual particles. We can 370 estimate that OM on average accounted for $63\pm23\%$ of the dry sulfate particle volume. 371 372 Our result shows that the OM volume increases following the particle size increase (Figure S6). 373

374

375 4. Discussion

4.1 Mixing mechanism of organic, soot, and sulfate

Lagrangian particle dispersion modeling using the FLEXPART-WRF 3.1 showed 377 that air masses arriving at the sampling site during our field measurement periods 378 were likely originated from the Greenland and North America (Figure 2). Previous 379 studies reported that air masses from North America or Greenland during the summer 380 contain higher concentration of black carbon, OM, and sulfate (Burkart et al., 2017; 381 382 Chang et al., 2011; Fu et al., 2008; Moore et al., 2011; Park et al., 2013). Indeed, there is strong emission intensity of OC and SO₂ around the Arctic area from emission 383 simulation as shown in Figure S2. However, Weinbruch et al. (2012) observed soot 384 particles when cruise ships were present in the area around Ny-Ålesund town. It is 385 possible that minor soot particles are sourced from the ship emissions and most of 386 them are transported from out of Arctic area in the free troposphere (Figure S2). 387

The sulfate core-OM shell structure observed in the Arctic summer atmosphere is 388 similar to those in the background or rural air in other places (Li et al., 2016; Moffet 389 390 et al., 2013). Based on the images from electron microscopies, we can infer that OM coating thickness in the Arctic air was comparable with them in rural places but 391 higher than them in urban places. During the transports, organic coatings on sulfates 392 were considered as the secondary organic aerosols and their masses increase 393 following particle aging and growth (Li et al., 2016; Moffet et al., 2013; Sierau et al., 394 2014). Figures 1 and 2 show that most of particles in the air masses transported long 395 distance from North American. The result indicates that these long-range 396 transportation of secondary sulfate particles have enough time to experience the 397 398 possible atmospheric heterogeneous reactions on particle surfaces or cloud processes 399 in the Arctic air. Similarly, Moffet et al. (2013) found that soot inclusions occurred in OM coating when OM coating on sulfates built up through photochemical activity and 400 pollution buildup the Sacramento urban plume aged. On the other hand, the 401 402 sulfate/OM particles with soot inclusions are probably formed in a similar way as 403 those found elsewhere (Li et al., 2016) – e.g., soot particles may have acted as nuclei for secondary sulfate or organic uptake during their transports (Riemer et al., 2009). 404

Similarly, besides the OM coating in the Arctic particles, Leck and Svensson (2015)
found some biogenic aerosols like gel-aggregate containing bacterium in ultrafine
particles. However, we didn't find any gel-like particles in the samples because our
sampler had very low efficiency for ultrafine particles.

TEM images show that most of the internally mixed sulfate particles display 409 sulfate core and OM coating on the substrate (Figures 4a and 5b, c). The sulfate and 410 OM separation in individual particles were defined by You et al. (2012) as 411 412 liquid-liquid phase separation (LLPS). Concerning the knowledges of the LLPS can better understand particle hygroscopicity, heterogeneous reactions of reactive gases on 413 particle surface, and organic aging (You et al., 2012). They also reported that the 414 LLPS can reflect the O:C ratio in the OM, which is roughly ≤ 0.5 . In this study, we 415 did observe the LLPS in almost all the fine sulfate particles, which indicates that the 416 secondary OM in the coating might be not highly aged. Therefore, we speculate that 417 the thick OM coatings were consistently built up during the long-range transport of 418 sulfate particles and part of secondary OM in the coating likely formed in Arctic area. 419 420 Indeed, some studies reported that there are various sources of organic precursors during the Arctic area, such as biogenic VOCs from ice melting and open water 421 (Dall Osto et al., 2017) and anthropogenic VOCs from shipping emissions in 422 summertime (Gilgen et al., 2018). The dependence of OM volume on particle size 423 (Figure S6) suggests that the suspended sulfate particles are initially important surface 424 for secondary OM formation. Moreover, the common OM coating on sulfate particles 425 indicates that secondary OM as the surfaces of fine particles might govern the 426 possible heterogeneous reactions between reactive gases and sulfate particles in the 427 428 Arctic air.

It should be noted that most of secondary OM not only occurred on the surfaces of sulfate particles but also its mass (mean mass at 63±23%) dominated in individual particles (Figure 7d). The OM dominating in individual particles can influence the IN and CCN activities of secondary sulfate particles (Lathem et al., 2013; Martin et al., 2011). For example, some studies found that an increase in organic mass fraction in particles of a certain size would lead to a suppression of the Arctic CCN activity (Leck and Svensson, 2015; Martin et al., 2011). Moreover, OM as particle surfaces
can significantly influence hygroscopicity and IN activity of sulfate particles (Wang et
al., 2012).

438

439 **4.2** Potential impact of OM on optical properties of sulfate-containing particles

The internal mixing of soot, sulfate, and OM can change optical properties of 440 individual particles in the atmosphere. Recent studies showed that BrC has been 441 442 detected in the OM in the polluted and clean air and even in upper troposphere (Laskin et al., 2015; Wang et al., 2018). Feng et al. (2013) further calculated the 443 contribution up to 19% of the optical absorption of the strongly absorbing BrC in 444 global simulations which is after the absorption BC aerosols. Although we didn't 445 directly measure the optical absorption and BrC in the Arctic atmosphere, various 446 colored OM (e.g. nitrated/polycyclic aromatics and phenols), referred as BrC, were 447 detected in the Arctic atmosphere in different seasons (Fu et al., 2008; 448 Wöhrnschimmel et al., 2013; Zangrando et al., 2013) and in surface ice or snowpack 449 450 (Browse et al., 2013; Doherty et al., 2013; Hegg et al., 2010). We also noticed that the ¹²C¹⁴N⁻ signal generally occurred in all analyzed OM coating in sulfate particles 451 (Figure 4e-f). Herrmann et al. (2007) considered that ¹²C¹⁴N⁻ from NanoSIMS 452 represents nitrogen-containing organic in the detected materials. In this study, 453 although we could not determine that all the organic materials in the OM coating were 454 nitrogen-containing OM, the NanoSIMS data as shown in Figure 4 indicated that the 455 OM coating more or less homogenously contained nitrogen-containing OM. As a 456 result, the nitrogen-containing OM indicates that the OM coating could contain 457 certain amounts of secondary BrC (Jiang et al., 2019; Laskin et al., 2015). 458

To understand how OM coating influence optical properties of sulfate particles, we assume three scenarios of OM coating as BrC: strongly absorbing (case 1), moderately absorbing (case 2) or non-absorbing OM (case 3) with a refractive index of 1.65-0.03i, 1.65-0.003i, and 1.65 at 550 nm according to *Feng et al.* (2013). Based on the size measurements shown in Figure 7d, we can calculate volume of sulfate and OM within each particle. We input volume of each component and the corresponding 465 refractive index into the Mie code and then calculated optical properties of individual sulfate particles in the samples. Based on optical data statistic of 575 particles, Figure 466 8a show that the OM coating is strongly absorbing BrC (referred to case Abs1), as by 467 Feng et al.(2013), the average absorption cross section (ACS) of individual particles is 468 estimated to be 2.67×10^{-14} m². This value is 8.30 times higher than the aerosol ACS 469 $(3.22 \times 10^{-15} \text{ m}^2)$ when assuming that the BrC is moderately absorbing (referred to case 470 Abs2, Figure 8a). However, the scattering cross section (SCS) of individual particles 471 472 only shows a small change (Figure 8b). Figure 8c also shows that the single scattering albedos (SSAs) of individual particles are 0.92, 0.99, and 1 when assuming the BrC as 473 strongly, moderately and non-absorbing (cases SSA1 to SSA3). These results suggest 474 whether we consider organic coating as BrC may have a significant influence on the 475 absorption properties of individual sulfate particles. 476

In this study, we expored the relationship between ACS of individual particles and particle diameters. Interestingly, Figure 8d shows that ACS of individual fine OM-coating sulfate particles increased following the increasing particle size. The result shows that the ACS can be enhanced following particle size growing and particle aging. In other word, OM-coating sulfate particles transported more longer distances and they might have stronger optical absorption in the Arctic air.

Current climate models estimated the radiative force of Arctic BC (Sand et al., 2013; Shindell, 2007; Winiger et al., 2017; Zanatta et al., 2018), but none specifically considered optical properties of Arctic BrC. Our study well revealed OM coating on sulfate particles and this detail microphysical complexity of aerosol particles will be useful to construct the atmospheric radiation and CCN/IN simulation in Arctic atmospheric models in the future.

489

490 **5** Summary

491 Different individual particle techniques, such as TEM/EDS, STEM, SEM, 492 NanoSIMS, and AFM, were applied to study S-rich, soot, and OM particles in the 493 Arctic air in summer. Sulfate particles accounted for approximately $29\pm7\%$ by 494 number of all analyzed particles in Arctic air. TEM and NanoSIMS commonly

observed OM coating and sulfate core individual sulfate particles, defined as the 495 LLSP. The common OM coating on sulfate particles indicates that secondary OM as 496 the surfaces of fine particles might govern the possible heterogeneous reactions 497 between reactive gases and sulfate particles in the Arctic air. Moreover, 20% of them 498 also contain small soot inclusions but they only appeared in organic coating, rather 499 than as the core mixed in sulfate. The mixing structure is totally different from the 500 previous findings that soot is internally mixed with sulfate instead of OM coating in 501 502 urban polluted air.

Size distribution of the secondary Arctic particles displayed a peak at 340 nm, 503 ranging from 100 nm to 2000 nm. The core particles, as sulfate or soot, had a peak at 504 240 nm and 120 nm, respectively. Furthermore, we can estimate that OM on average 505 accounted for 63±23% of the dry NSS-particle volume. Based on microscopic 506 measurements of individual particles, we not only built up one core-shell model but 507 also quantify volume of OM and sulfate in individual particles. The Mie code was 508 used to calculate optical properties of internally mixed sulfate/OM particles when we 509 510 considered OM as non-absorbing, moderately absorbing BrC, and strongly absorbing BrC. We found that the aerosol ACS is 8.30 times higher than the BrC as moderately 511 absorbing. We concluded that whether we consider organic coating as BrC may have a 512 significant influence on the absorption properties of individual particles in the Arctic 513 air. Moreover, individual fine OM-coating sulfate particles increased following the 514 increasing particle size. Therefore, we proposed that further studies should focus on 515 the BrC in Arctic aerosols: What mass concentrations of BrC are in fine particles? 516 What kinds of BrC are in fine particles? The optical mass absorption of BrC in fine 517 particles should be investigated? These results can be used to evaluate how BrC 518 519 aerosols influence the Arctic climate.

520

521	Author Contributions: WL and ZS designed the study. YZ and XS collected aerosol
522	particles. WL, HY, and JZ contributed laboratory experiments and data analysis. HY
523	and WL performed optical calculation and wrote part of first draft. PT and MD
524	provided the online measurement data of new particle formation and growth. JS and
525	XZ coordinated the field campaign. All authors commented and edited the paper.
526	
527	Competing interests: The authors declare no competing financial interests
528	
529	Acknowledgments We thank Boris Quennehen to provide data from the
530	FLEXPART-WRF. This work was funded by National Natural Science Foundation of
531	China (41622504, 41575116, 31700475) and the Hundred Talents Program in

532 Zhejiang University, Z.S. acknowledges funding from NERC (NE/S00579X/1).

References

- Abbatt, J.P.D., Leaitch, W.R., Aliabadi, A.A., Bertram, A.K., Blanchet, J.P., Boivin-Rioux, A., Bozem, H., Burkart, J., Chang, R.Y.W., Charette, J., Chaubey, J.P., Christensen, R.J., Cirisan, A., Collins, D.B., Croft, B., Dionne, J., Evans, G.J., Fletcher, C.G., Gal í M., Ghahremaninezhad, R., Girard, E., Gong, W., Gosselin, M., Gourdal, M., Hanna, S.J., Hayashida, H., Herber, A.B., Hesaraki, S., Hoor, P., Huang, L., Hussherr, R., Irish, V.E., Keita, S.A., Kodros, J.K., Kölner, F., Kolonjari, F., Kunkel, D., Ladino, L.A., Law, K., Levasseur, M., Libois, Q., Liggio, J., Lizotte, M., Macdonald, K.M., Mahmood, R., Martin, R.V., Mason, R.H., Miller, L.A., Moravek, A., Mortenson, E., Mungall, E.L., Murphy, J.G., Namazi, M., Norman, A.L., O'Neill, N.T., Pierce, J.R., Russell, L.M., Schneider, J., Schulz, H., Sharma, S., Si, M., Staebler, R.M., Steiner, N.S., Thomas, J.L., von Salzen, K., Wentzell, J.J.B., Willis, M.D., Wentworth, G.R., Xu, J.W., Yakobi-Hancock, J.D.: Overview paper: New insights into aerosol and climate in the Arctic, Atmos. Chem. Phys., 19 (4), 2527-2560, 2019.
- Alexander, D.T.L., Crozier, P.A., Anderson, J.R.: Brown Carbon Spheres in East Asian Outflow and Their Optical Properties, Science, 321 (5890), 833-836, 2008.
- Andreae, M.O., Gelencser, A.: Black carbon or brown carbon? The nature of light-absorbing carbonaceous aerosols, Atmos. Chem. Phys., 6 (10), 3131-3148, 2006.
- Behrenfeldt, U., Krejci, R., Ström, J., Stohl, A.: Chemical properties of Arctic aerosol particles collected at the Zeppelin station during the aerosol transition period in May and June of 2004, Tellus B, 60 (3), 405-415, 2008.
- Bohren, C.F., Huffman, D.R., 1983. Absorption and scattering of light by small particles. John Wiley & Sons, Inc., New York, USA.
- Bond, T.C., Doherty, S.J., Fahey, D.W., Forster, P.M., Berntsen, T., DeAngelo, B.J., Flanner, M.G., Ghan, S., Kärcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P.K., Sarofim, M.C., Schultz, M.G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S., Bellouin, N., Guttikunda, S.K., Hopke, P.K., Jacobson, M.Z., Kaiser, J.W., Klimont, Z., Lohmann, U., Schwarz, J.P., Shindell, D., Storelvmo, T., Warren, S.G., Zender, C.S.: Bounding the role of black carbon in the climate system: A scientific assessment, J. Geophys. Res., 118 (11), 5380-5552, 2013.
- Brioude, J., Arnold, D., Stohl, A., Cassiani, M., Morton, D., Seibert, P., Angevine, W., Evan, S., Dingwell, A., Fast, J.D., Easter, R.C., Pisso, I., Burkhart, J., Wotawa, G.: The Lagrangian particle dispersion model FLEXPART-WRF version 3.1, Geosci. Model Dev., 6 (6), 1889-1904, 2013.
- Brock, C.A., Cozic, J., Bahreini, R., Froyd, K.D., Middlebrook, A.M., McComiskey, A., Brioude, J., Cooper, O.R., Stohl, A., Aikin, K.C., de Gouw, J.A., Fahey, D.W., Ferrare, R.A., Gao, R.S., Gore, W., Holloway, J.S., Hübler, G., Jefferson, A., Lack, D.A., Lance, S., Moore, R.H., Murphy, D.M., Nenes, A., Novelli, P.C., Nowak, J.B., Ogren, J.A., Peischl, J., Pierce, R.B., Pilewskie, P., Quinn, P.K., Ryerson, T.B., Schmidt, K.S., Schwarz, J.P., Sodemann, H., Spackman, J.R., Stark, H., Thomson, D.S., Thornberry, T., Veres, P., Watts, L.A., Warneke, C., Wollny, A.G.: Characteristics, sources, and transport of aerosols measured in spring 2008 during the aerosol, radiation, and cloud processes affecting Arctic Climate (ARCPAC) Project, Atmos. Chem. Phys., 11 (6), 2423-2453, 2011.
- Browse, J., Carslaw, K.S., Schmidt, A., Corbett, J.J.: Impact of future Arctic shipping on high-latitude black carbon deposition, Geophys. Res. Lett., 40 (16), 4459-4463, 2013.
- Burkart, J., Willis, M.D., Bozem, H., Thomas, J.L., Law, K., Hoor, P., Aliabadi, A.A., Köllner, F., Schneider, J., Herber, A., Abbatt, J.P.D., Leaitch, W.R.: Summertime observations of elevated levels of ultrafine particles in the high Arctic marine boundary layer, Atmos. Chem. Phys., 17 (8),

5515-5535, 2017.

- Buseck, P.R., Posfai, M.: Airborne minerals and related aerosol particles: Effects on climate and the environment, P. Natl. Acad. Sci. USA, 96 (7), 3372-3379, 1999.
- Chang, R.Y.W., Leck, C., Graus, M., Müller, M., Paatero, J., Burkhart, J.F., Stohl, A., Orr, L.H., Hayden, K., Li, S.M., Hansel, A., Tjernström, M., Leaitch, W.R., Abbatt, J.P.D.: Aerosol composition and sources in the central Arctic Ocean during ASCOS, Atmos. Chem. Phys., 11 (20), 10619-10636, 2011.
- Chi, J.W., Li, W.J., Zhang, D.Z., Zhang, J.C., Lin, Y.T., Shen, X.J., Sun, J.Y., Chen, J.M., Zhang, X.Y., Zhang, Y.M., Wang, W.X.: Sea salt aerosols as a reactive surface for inorganic and organic acidic gases in the Arctic troposphere, Atmos. Chem. Phys., 15 (19), 11341-11353, 2015.
- Dagsson-Waldhauserova, P., Arnalds, O., Olafsson, H.: Long-term frequency and characteristics of dust storm events in Northeast Iceland (1949–2011), Atmos. Environ., 77 (0), 117-127, 2013.
- Dall Osto, M., Beddows, D.C.S., Tunved, P., Krejci, R., Ström, J., Hansson, H.C., Yoon, Y.J., Park, K.-T., Becagli, S., Udisti, R., Onasch, T., O Dowd, C.D., Simó, R., Harrison, R.M.: Arctic sea ice melt leads to atmospheric new particle formation, Sci. Rep., 7 (1), 3318, 2017.
- Doherty, S.J., Grenfell, T.C., Forsström, S., Hegg, D.L., Brandt, R.E., Warren, S.G.: Observed vertical redistribution of black carbon and other insoluble light-absorbing particles in melting snow, J. Geophys. Res., 118 (11), 5553-5569, 2013.
- Feng, Y., Ramanathan, V., Kotamarthi, V.R.: Brown carbon: a significant atmospheric absorber of solar radiation?, Atmos. Chem. Phys., 13 (17), 8607-8621, 2013.
- Fu, P., Kawamura, K., Barrie, L.A.: Photochemical and Other Sources of Organic Compounds in the Canadian High Arctic Aerosol Pollution during Winter–Spring, Environ. Sci. Technol., 43 (2), 286-292, 2008.
- Geng, H., Ryu, J., Jung, H.-J., Chung, H., Ahn, K.-H., Ro, C.-U.: Single-Particle Characterization of Summertime Arctic Aerosols Collected at Ny-Alesund, Svalbard, Environ. Sci. Technol., 44 (7), 2348-2353, 2010.
- Ghosal, S., Weber, P.K., Laskin, A.: Spatially resolved chemical imaging of individual atmospheric particles using nanoscale imaging mass spectrometry: insight into particle origin and chemistry, Analytical Methods, 6 (8), 2444-2451, 2014.
- Gilgen, A., Huang, W.T.K., Ickes, L., Neubauer, D., Lohmann, U.: How important are future marine and shipping aerosol emissions in a warming Arctic summer and autumn?, Atmos. Chem. Phys., 18 (14), 10521-10555, 2018.
- Hansen, J., Nazarenko, L.: Soot climate forcing via snow and ice albedos, P. Natl. Acad. Sci. USA, 101 (2), 423-428, 2004.
- Hara, K., Yamagata, S., Yamanouchi, T., Sato, K., Herber, A., Iwasaka, Y., Nagatani, M., Nakata, H.: Mixing states of individual aerosol particles in spring Arctic troposphere during ASTAR 2000 campaign, J. Geophys. Res., 108 (D7), 2003.
- Hegg, D.A., Warren, S.G., Grenfell, T.C., Sarah, J.D., Clarke, A.D.: Sources of light-absorbing aerosol in arctic snow and their seasonal variation, Atmos. Chem. Phys., 10 (22), 10923-10938, 2010.
- Herrmann, A.M., Ritz, K., Nunan, N., Clode, P.L., Pett-Ridge, J., Kilburn, M.R., Murphy, D.V., O'Donnell, A.G., Stockdale, E.A.: Nano-scale secondary ion mass spectrometry — A new analytical tool in biogeochemistry and soil ecology: A review article, Soil Biol. Biochem., 39 (8), 1835-1850, 2007.
- IPCC (2013), Clouds and Aerosols, in Climate Change 2013: The Physical Science Basis. Contribution

of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, 571-657 pp, Cambridge, U.K. and New York, NY.

- Iwasaka, Y., Minoura, H., Nagaya, K.: The transport and spatial scale of Asian dust-storm clouds: A case study of the dust-storm event of April 1979, Tellus, Ser. B, 35, 189-196, 1983.
- Iziomon, M.G., Lohmann, U., Quinn, P.K.: Summertime pollution events in the Arctic and potential implications, J. Geophys. Res., 111 (D12), D12206, 2006.
- Jacob, D.J., Crawford, J.H., Maring, H., Clarke, A.D., Dibb, J.E., Emmons, L.K., Ferrare, R.A., Hostetler, C.A., Russell, P.B., Singh, H.B., Thompson, A.M., Shaw, G.E., McCauley, E., Pederson, J.R., Fisher, J.A.: The Arctic Research of the Composition of the Troposphere from Aircraft and Satellites (ARCTAS) mission: design, execution, and first results, Atmos. Chem. Phys., 10 (11), 5191-5212, 2010.
- Jiang, H., Frie, A.L., Lavi, A., Chen, J.Y., Zhang, H., Bahreini, R., Lin, Y.-H.: Brown Carbon Formation from Nighttime Chemistry of Unsaturated Heterocyclic Volatile Organic Compounds, Environ. Sci. Techn. Lett., DOI: 10.1021/acs.estlett.1029b00017, 2019.
- Karl, M., Leck, C., Coz, E., Heintzenberg, J.: Marine nanogels as a source of atmospheric nanoparticles in the high Arctic, Geophys. Res. Lett., 40 (14), 3738-3743, 2013.
- Kirpes, R.M., Bondy, A.L., Bonanno, D., Moffet, R.C., Wang, B., Laskin, A., Ault, A.P., Pratt, K.A.: Secondary sulfate is internally mixed with sea spray aerosol and organic aerosol in the winter Arctic, Atmos. Chem. Phys., 18 (6), 3937-3949, 2018.
- Koch, D., Hansen, J.: Distant origins of Arctic black carbon: A Goddard Institute for Space Studies ModelE experiment, J. Geophys. Res., 110 (D4), D04204, 2005.
- Lack, D.A., Langridge, J.M., Bahreini, R., Cappa, C.D., Middlebrook, A.M., Schwarz, J.P.: Brown carbon and internal mixing in biomass burning particles, P. Natl. Acad. Sci. USA, 109 (37), 14802-14807, 2012.
- Laskin, A., Laskin, J., Nizkorodov, S.A.: Chemistry of Atmospheric Brown Carbon, Chem. Rev., 115 (10), 4355-4382, 2015.
- Laskina, O., Morris, H.S., Grandquist, J.R., Estillore, A.D., Stone, E.A., Grassian, V.H., Tivanski, A.V.: Substrate-Deposited Sea Spray Aerosol Particles: Influence of Analytical Method, Substrate, and Storage Conditions on Particle Size, Phase, and Morphology, Environ. Sci. Tech., 49 (22), 13447-13453, 2015.
- Lathem, T.L., Beyersdorf, A.J., Thornhill, K.L., Winstead, E.L., Cubison, M.J., Hecobian, A., Jimenez, J.L., Weber, R.J., Anderson, B.E., Nenes, A.: Analysis of CCN activity of Arctic aerosol and Canadian biomass burning during summer 2008, Atmos. Chem. Phys., 13 (5), 2735-2756, 2013.
- Law, K.S., Stohl, A.: Arctic Air Pollution: Origins and Impacts, Science, 315 (5818), 1537-1540, 2007.
- Leck, C., Bigg, E.K.: Comparison of sources and nature of the tropical aerosol with the summer high Arctic aerosol, Tellus, Ser. B, 60 (1), 118-126, 2008.
- Leck, C., Svensson, E.: Importance of aerosol composition and mixing state for cloud droplet activation over the Arctic pack ice in summer, Atmos. Chem. Phys., 15 (5), 2545-2568, 2015.
- Li, W., Sun, J., Xu, L., Shi, Z., Riemer, N., Sun, Y., Fu, P., Zhang, J., Lin, Y., Wang, X., Shao, L., Chen, J., Zhang, X., Wang, Z., Wang, W.: A conceptual framework for mixing structures in individual aerosol particles, J. Geophys. Res., 121 (22), 13,784-713,798, 2016.
- Li, W., Xu, L., Liu, X., Zhang, J., Lin, Y., Yao, X., Gao, H., Zhang, D., Chen, J., Wang, W., Harrison, R.M., Zhang, X., Shao, L., Fu, P., Nenes, A., Shi, Z.: Air pollution–aerosol interactions produce more bioavailable iron for ocean ecosystems, Sci. Adv., 3 (3), e1601749, 2017.

- Maahn, M., de Boer, G., Creamean, J.M., Feingold, G., McFarquhar, G.M., Wu, W., Mei, F.: The observed influence of local anthropogenic pollution on northern Alaskan cloud properties, Atmos. Chem. Phys., 17 (23), 14709-14726, 2017.
- Martin, M., Chang, R.Y.W., Sierau, B., Sjogren, S., Swietlicki, E., Abbatt, J.P.D., Leck, C., Lohmann, U.: Cloud condensation nuclei closure study on summer arctic aerosol, Atmos. Chem. Phys., 11 (22), 11335-11350, 2011.
- Moffet, R.C., Rödel, T.C., Kelly, S.T., Yu, X.Y., Carroll, G.T., Fast, J., Zaveri, R.A., Laskin, A., Gilles, M.K.: Spectro-microscopic measurements of carbonaceous aerosol aging in Central California, Atmos. Chem. Phys., 13 (20), 10445-10459, 2013.
- Moore, R.H., Bahreini, R., Brock, C.A., Froyd, K.D., Cozic, J., Holloway, J.S., Middlebrook, A.M., Murphy, D.M., Nenes, A.: Hygroscopicity and composition of Alaskan Arctic CCN during April 2008, Atmos. Chem. Phys., 11 (22), 11807-11825, 2011.
- Moroni, B., Cappelletti, D., Crocchianti, S., Becagli, S., Caiazzo, L., Traversi, R., Udisti, R., Mazzola, M., Markowicz, K., Ritter, C., Zielinski, T.: Morphochemical characteristics and mixing state of long range transported wildfire particles at Ny-Ålesund (Svalbard Islands), Atmos. Environ., 156, 135-145, 2017.
- Park, K., Kim, G., Kim, J.-s., Yoon, Y.-J., Cho, H.-j., Ström, J.: Mixing State of Size-Selected Submicrometer Particles in the Arctic in May and September 2012, Environ. Sci. Technol., 48 (2), 909-919, 2013.
- Peters, G.P., Nilssen, T.B., Lindholt, L., Eide, M.S., Glomsrød, S., Eide, L.I., Fuglestvedt, J.S.: Future emissions from shipping and petroleum activities in the Arctic, Atmos. Chem. Phys., 11 (11), 5305-5320, 2011.
- Qi, L., Li, Q., Li, Y., He, C.: Factors controlling black carbon distribution in the Arctic, Atmos. Chem. Phys., 17 (2), 1037-1059, 2017.
- Quinn, P.K., Shaw, G., Andrews, E., Dutton, E.G., Ruoho-Airola, T., Gong, S.L.: Arctic haze: current trends and knowledge gaps, Tellus B, 59 (1), 99-114, 2007.
- Raatikainen, T., Brus, D., Hyvärinen, A.P., Svensson, J., Asmi, E., Lihavainen, H.: Black carbon concentrations and mixing state in the Finnish Arctic, Atmos. Chem. Phys., 15 (17), 10057-10070, 2015.
- Riemer, N., West, M., Zaveri, R.A., Easter, R.C.: Simulating the evolution of soot mixing state with a particle resolved aerosol model, J. Geophys. Res., 114, doi:10.1029/2008JD011073, 2009.
- Saleh, R., Hennigan, C.J., McMeeking, G.R., Chuang, W.K., Robinson, E.S., Coe, H., Donahue, N.M., Robinson, A.L.: Absorptivity of brown carbon in fresh and photo-chemically aged biomass-burning emissions, Atmos. Chem. Phys., 13 (15), 7683-7693, 2013.
- Samset, B.H., Myhre, G., Herber, A., Kondo, Y., Li, S.M., Moteki, N., Koike, M., Oshima, N., Schwarz, J.P., Balkanski, Y., Bauer, S.E., Bellouin, N., Berntsen, T.K., Bian, H., Chin, M., Diehl, T., Easter, R.C., Ghan, S.J., Iversen, T., Kirkev åg, A., Lamarque, J.F., Lin, G., Liu, X., Penner, J.E., Schulz, M., Seland, Ø., Skeie, R.B., Stier, P., Takemura, T., Tsigaridis, K., Zhang, K.: Modelled black carbon radiative forcing and atmospheric lifetime in AeroCom Phase II constrained by aircraft observations, Atmos. Chem. Phys., 14 (22), 12465-12477, 2014.
- Sand, M., Berntsen, T.K., Kay, J.E., Lamarque, J.F., Seland, Ø., Kirkev åg, A.: The Arctic response to remote and local forcing of black carbon, Atmos. Chem. Phys., 13 (1), 211-224, 2013.
- Seinfeld, J., Pandis, S., 2006. Atmospheric Chemistry and Physics: From air pollution to climate change (2nd ed.). 1-1203 pp., John Wiley & Son, Inc., Hoboken, New Jersey.

- Shindell, D.: Local and remote contributions to Arctic warming, Geophys. Res. Lett., 34 (14), L14704, 2007.
- Sierau, B., Chang, R.Y.W., Leck, C., Paatero, J., Lohmann, U.: Single-particle characterization of the high-Arctic summertime aerosol, Atmos. Chem. Phys., 14 (14), 7409-7430, 2014.
- Updyke, K.M., Nguyen, T.B., Nizkorodov, S.A.: Formation of brown carbon via reactions of ammonia with secondary organic aerosols from biogenic and anthropogenic precursors, Atmos. Environ., 63 (0), 22-31, 2012.
- Wöhrnschimmel, H., MacLeod, M., Hungerbuhler, K.: Emissions, Fate and Transport of Persistent Organic Pollutants to the Arctic in a Changing Global Climate, Environ. Sci. Technol., 47 (5), 2323-2330, 2013.
- Wang, B., Laskin, A., Roedel, T., Gilles, M.K., Moffet, R.C., Tivanski, A.V., Knopf, D.A.: Heterogeneous ice nucleation and water uptake by field-collected atmospheric particles below 273 K, J. Geophys. Res., 117, 2012.
- Wang, X., Heald, C.L., Liu, J., Weber, R.J., Campuzano-Jost, P., Jimenez, J.L., Schwarz, J.P., Perring, A.E.: Exploring the observational constraints on the simulation of brown carbon, Atmos. Chem. Phys., 18 (2), 635-653, 2018.
- Weinbruch, S., Wiesemann, D., Ebert, M., Schütze, K., Kallenborn, R., Ström, J.: Chemical composition and sources of aerosol particles at Zeppelin Mountain (Ny Ålesund, Svalbard): An electron microscopy study, Atmos. Environ., 49 (0), 142-150, 2012.
- Winiger, P., Andersson, A., Eckhardt, S., Stohl, A., Gustafsson, Ö.: The sources of atmospheric black carbon at a European gateway to the Arctic, Nat. Commun., 7, 12776, 2016.
- Winiger, P., Andersson, A., Eckhardt, S., Stohl, A., Semiletov, I.P., Dudarev, O.V., Charkin, A., Shakhova, N., Klimont, Z., Heyes, C., Gustafsson, Ö.: Siberian Arctic black carbon sources constrained by model and observation, P. Natl. Acad. Sci. USA, 114 (7), E1054-E1061, 2017.
- Xu, J.W., Martin, R.V., Morrow, A., Sharma, S., Huang, L., Leaitch, W.R., Burkart, J., Schulz, H., Zanatta, M., Willis, M.D., Henze, D.K., Lee, C.J., Herber, A.B., Abbatt, J.P.D.: Source attribution of Arctic black carbon constrained by aircraft and surface measurements, Atmos. Chem. Phys., 17 (19), 11971-11989, 2017.
- Yang, Y., Wang, H., Smith, S.J., Easter, R.C., Rasch, P.J.: Sulfate Aerosol in the Arctic: Source Attribution and Radiative Forcing, J. Geophys. Res., 123 (3), 1899-1918, 2018.
- You, Y., Renbaum-Wolff, L., Carreras-Sospedra, M., Hanna, S.J., Hiranuma, N., Kamal, S., Smith, M.L., Zhang, X., Weber, R.J., Shilling, J.E., Dabdub, D., Martin, S.T., Bertram, A.K.: Images reveal that atmospheric particles can undergo liquid–liquid phase separations, P. Natl. Acad. Sci. USA, 109 (33), 13188-13193, 2012.
- Zanatta, M., Laj, P., Gysel, M., Baltensperger, U., Vratolis, S., Eleftheriadis, K., Kondo, Y., Dubuisson, P., Winiarek, V., Kazadzis, S., Tunved, P., Jacobi, H.W.: Effects of mixing state on optical and radiative properties of black carbon in the European Arctic, Atmos. Chem. Phys., 18 (19), 14037-14057, 2018.
- Zangrando, R., Barbaro, E., Zennaro, P., Rossi, S., Kehrwald, N.M., Gabrieli, J., Barbante, C., Gambaro, A.: Molecular Markers of Biomass Burning in Arctic Aerosols, Environ. Sci. Technol., 47 (15), 8565-8574, 2013.

Figure Captions

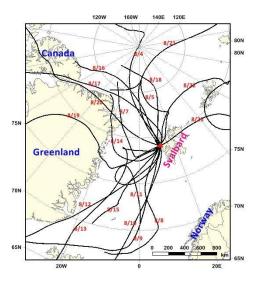


Figure 1 72 h back trajectories of air masses at 500m over Arctic Yellow River Station in Svalbard during 3–26 August 2012, and arriving time was set according to the sampling time

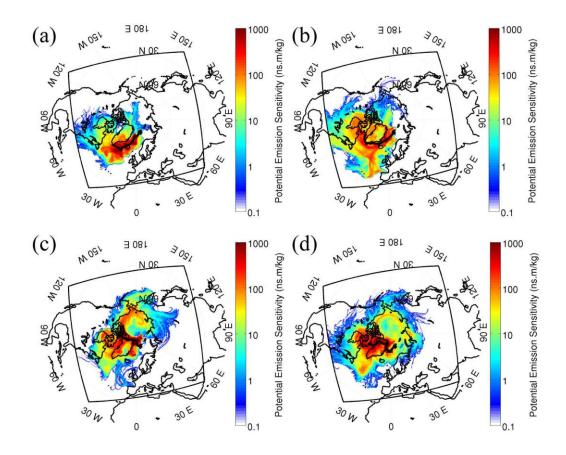


Figure 2 FLEXPART-WRF PES on August 11, 12, 14, and 15, 2012. Black square is showing the WRF domain used to initiate the FLEXPART-WRF simulation.

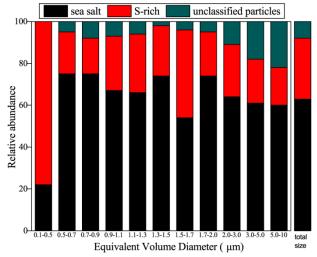


Figure 3 Morphology and relative abundances of typical individual aerosol particles in the 21 analyzed samples.

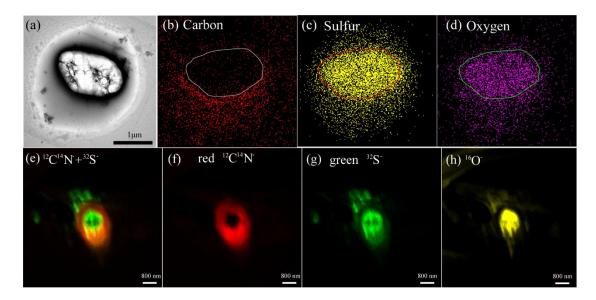


Figure 4 TEM Observations of a secondary particle and NanoSIMS intensity threshold maps of an aerosol particle with sulfate core and OM coating. (a) Bright-field TEM image of an internally mixed particle; (b) elemental carbon (c) sulfur and (d) oxygen maps of the internally mixed particle shown in 1(a); (e) Overlay of ${}^{12}C{}^{14}N{}^{-}$ and ${}^{32}S{}^{-}$ ion maps in an internally mixed particle; (f) CN ${}^{-}$ map (g) S ${}^{-}$ (h) O ${}^{-}$ secondary ion maps. Ion maps with a set of aerosol particles were shown in Figure S1.

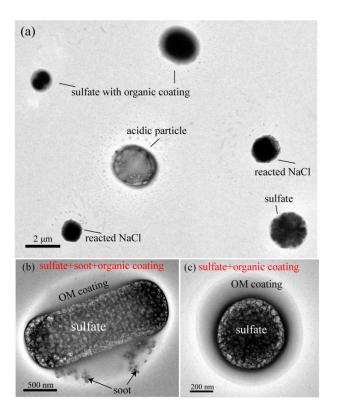


Figure 5 TEM images of individual particles containing sulfate, OM, and soot. (a) Low magnification TEM image showing sulfates, sulfate with OM coating, and reacted NaCl particles. (b) an internally mixed particle of sulfate and soot with OM coating (c) a particle with sulfate core and OM coating.

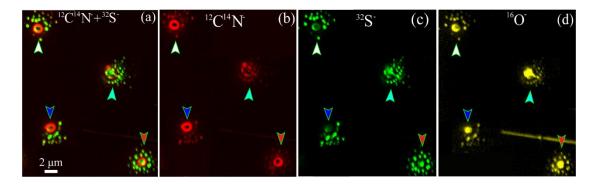


Figure 6 NanoSIMS intensity threshold maps of individual aerosol particles surrounded by satellite particles. (e) Overlay of ${}^{12}C{}^{14}N^{-}$ and ${}^{32}S^{-}$ ion maps of individual particles. (f) CN^{-} (g) S^{-} (h) O^{-} maps. Four particles were indicated by white, pink, blue, and red arrows.

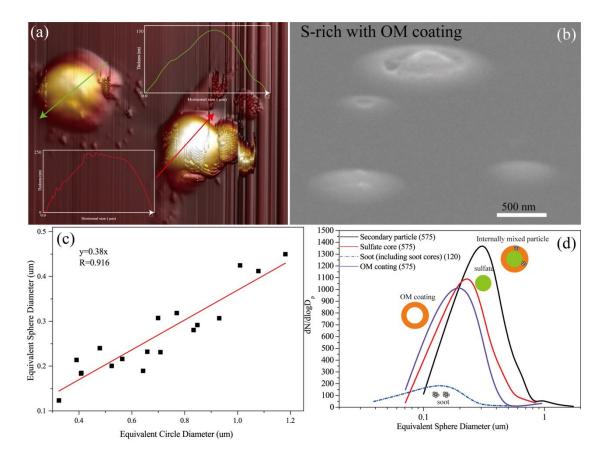


Figure 7 Secondary particles on the substrate. (a) 3-D AFM image of secondary sulfate particles. The colorful arrows represent particles surface properites of the particle section. (b) SEM image of S-rich with OM coating obtained from 75° tilt of the SEM specimen stage (c) The near linear relationships between ECD and ESD based on S-rich particles with thick OM coating by Atomic force microscopy. (d) Size distribution of individual particle with OM coating and sulfate cores based on the estimated ESD diameter from TEM image. Sizes of soot particles are equal to the equivalent circle diameter.

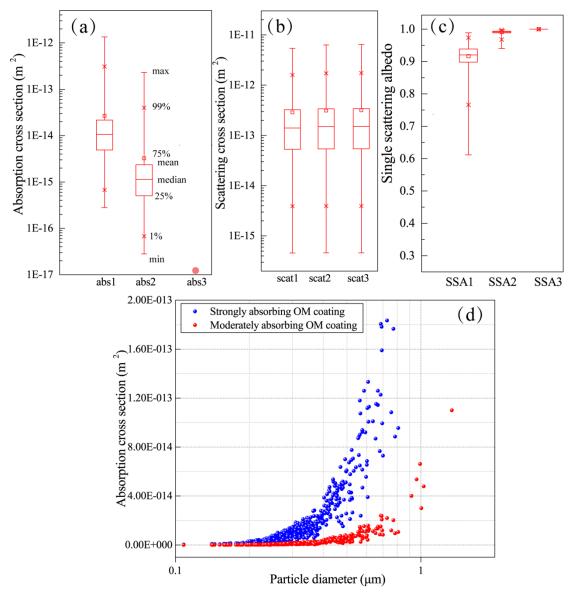


Figure 8 Optical properties of Box-and-whisker plots showing optical parameters of all analysed particles assuming sulfate core and BrC shell (not considering soot cores in the particles). (a) Scattering cross section (b) Absorption cross section (c) Single scattering albedo. Top to bottom makers in the box-and-whisker represent max, 99%, 75%, mean, median, 25%, 1%, min values. (d) Absorption cross section along with particle diameter assuming strongly absorbing BrC and Moderate absorbing BrC as the particle OM coating.