



1 **Measuring Light Absorption by Organic Aerosols: Correction Factors for Solvent**

2 **Extraction-Based Photometry Techniques**

3 Nishit Shetty¹, Apoorva Pandey¹, Stephen Baker², Wei Min Hao², Rajan K. Chakrabarty^{1,3}

4 ¹Center for Aerosol Science and Engineering, Department of Energy, Environmental and Chemical
5 Engineering, Washington University in St. Louis, St. Louis, MO 63130, USA

6 ²USDA Forest Service, Rocky Mountain Research Station, Fire Sciences Laboratory, Missoula, Montana,
7 USA

8 ³McDonnell Center for the Space Sciences, Washington University in St. Louis, St. Louis, MO 63130,
9 USA

10 Correspondence to: Rajan K. Chakrabarty (chakrabarty@wustl.edu)

11 **Abstract**

12 Recent studies have shown that organic aerosol (OA) could have a non-trivial role in
13 atmospheric light absorption at shorter visible wavelengths. Good estimates of OA absorption
14 are therefore necessary to accurately calculate radiative forcing due to these aerosols in climate
15 models. One of the common techniques used to measure OA light absorption is the solvent
16 extraction technique from filter samples which involves the use of a spectrophotometer to
17 measure bulk absorbance of the solvent-soluble organic fraction of particulate matter. Measured
18 bulk absorbance is subsequently converted to particle-phase absorption coefficient using
19 correction factors. The appropriate correction factors to use for performing this conversion under
20 varying scenarios of organic carbon (OC) to total carbon (TC) mass ratios has been an
21 unexplored area of research. The conventional view is to apply a correction factor of 2 for water-
22 extracted OA based on Mie calculations.

23 Here, we performed a comprehensive laboratory study involving three solvents (water, methanol,
24 and acetone) to investigate the corrections factors for converting from bulk-to-particle phase



25 absorption coefficients ($b_{\text{abs,OA}}/b_{\text{abs,bulk}}$) for primary OA emitted from biomass burning. We
26 parametrized these correction factors as a function of OC/TC mass ratio and single scattering
27 albedo (SSA). We observed these correction factors to be a function of the OC/TC ratio of the
28 aerosol, and that the conventionally used correction factor of 2 for water-extracted OA could
29 severely underpredict OA absorption at high EC mass fractions. We recommend using
30 $b_{\text{abs,OA}}/b_{\text{abs,bulk}}$ values between 2 and 11 for water extracts and values between 1 and 4 for
31 methanol extracts based on OC/TC ratios, for EC mass fractions less than 0.25. Furthermore, a
32 linear correlation between SSA and OC/TC ratio was also established. Finally, from the
33 spectroscopic data, we analyzed the differences in Absorption Ångström Exponents (AÅE)
34 obtained from bulk- and particulate-phase measurements. We noted that AÅE from bulk
35 measurements deviate significantly from their OA counterparts.

36 **1 Introduction**

37 Carbonaceous aerosols constitute a major short-lived climate pollutant, and even though they
38 have been studied extensively in recent years, estimates of their contribution to shortwave
39 radiative forcing remains highly uncertain (IPCC, 2013). Based on their thermal-refractory
40 properties, carbonaceous aerosols are categorized as elemental carbon (EC) or organic carbon
41 (OC) (Chow et al., 2007b; Bond et al., 2013), and the sum of OC and EC is referred to as total
42 carbon (TC). When defined optically, the refractory EC component is approximately referred to
43 as black carbon (BC) (Chow et al. 2007b; Bond et al., 2013); BC aerosol constitute the strongest
44 of the light absorbing aerosol components in the atmosphere (Andreae and Gelencsér 2006;
45 Ramanathan and Carmichael 2008; IPCC, 2013). While BC absorbs strongly in the visible
46 spectrum, the contribution of OC towards absorption has largely been neglected, even though
47 many studies have demonstrated significant OC absorption at lower visible wavelengths (Yang et



48 al., 2009; Chen and Bond, 2010; Chakrabarty et al., 2010; Chen et al., 2018). The atmospheric
49 mass of OC can be 3-12 times larger than that of BC (Husain et al., 2007; Zhang et al., 2008)
50 which warrants its inclusion as an atmospheric light absorber. Only recently have a few global
51 modeling studies incorporated radiative forcing by organic aerosol (OA) absorption (Chung et
52 al., 2012; Feng et al., 2013; Lin et al., 2014). Thus, having accurate estimates for OA absorption
53 is necessary to help improve climate models.

54 A convenient and prevalent methodology of measuring OA absorption is based on collecting
55 aerosol particles on a filter substrate followed by extracting the organic compounds into a
56 solvent. This is a good analytical method used in many studies as it excludes any interference
57 from EC and provides only the OC absorption spectra (Mo et al., 2017; Chen and Bond, 2010;
58 Liu et al., 2013). The absorbance of the organic chromophores in the solvent extract is measured
59 using an ultraviolet-visible (UV-Vis) spectrophotometer and the imaginary part of the complex
60 refractive index of the solvent is calculated using the measured absorbance and extract
61 concentrations. The absorbance values can be converted to corresponding bulk phase absorption
62 coefficients ($b_{\text{abs,bulk}}$), and the imaginary index along with an assumed number size distribution
63 can be used as inputs to Mie theory for calculating the absorption coefficient for the dissolved
64 OC particles. Past studies have suggested that $b_{\text{abs,bulk}}$ is not representative of the corresponding
65 particulate-phase organic aerosol absorption coefficient ($b_{\text{abs,OA}}$) (Liu et al., 2013; Moosmüller et
66 al., 2011). The situation thus warrants appropriate correction factors be applied to account for the
67 size-dependent absorption properties of OA when estimating $b_{\text{abs,OA}}$ using the solvent-extract
68 methodology. Using absorption coefficients determined with Mie Theory, previous studies have
69 determined a narrow range of solvent-dependent correction factors from 2 for water extracts to
70 1.8 for methanol extracts, all corresponding to a mean particle diameter of 0.5 μm (Liu et al.,



71 2013; Liu et al., 2016; Liu et al., 2014; Washenfelder et al., 2015). These correction factors have
72 been deemed relatively independent of particle size as long as the diameters do not get much
73 smaller than the wavelength of light, in which case their values fall to a range of 0.69 - 0.75 (Sun
74 et al., 2007). Further, these factors were calculated assuming homogeneous composition and
75 external mixing states of the OC aerosol in addition to fixed values of the real part of the
76 refractive index and effective density. No attempts have been made to quantify the variation in
77 these correction factors with varying aerosol intrinsic properties, such as the single scattering
78 albedo, and EC/OC ratios, even though these properties influence the types and fractions of
79 organics extracted by a given solvent (Zhang et al., 2013; Saleh et al., 2014) thereby affecting
80 $b_{\text{abs,bulk}}$.

81 In-situ measurement of particulate-phase absorption coefficient is commonly and accurately
82 accomplished using a photoacoustic spectrometer (PAS) (Lack et al., 2006; Arnott et al., 2005;
83 Arnott et al., 2003). However, a single-wavelength PAS cannot distinguish between absorption
84 by OC and BC aerosol and it typically measures the total particle-phase absorption coefficient
85 ($b_{\text{abs,tot}}$) of the aerosol population in the cell (Moosmüller et al., 2009). One must make use of a
86 multi-wavelength PAS using which $b_{\text{abs,OA}}$ could be separated out from that of BC, based on the
87 difference in their Absorption Ångström Exponent (AÅE) (Washenfelder et al., 2015; Arola et
88 al., 2011; Kirchstetter and Thatcher, 2012). The AÅE for pure BC is well-constrained at 1 in the
89 visible and near-infrared wavelengths (Moosmüller et al., 2009). The value of $b_{\text{abs,OA}}$ is then the
90 difference between $b_{\text{abs,tot}}$ and the BC absorption coefficient. A correction factor ($b_{\text{abs,OA}}/b_{\text{abs,bulk}}$)
91 can thus be determined, provided simultaneous measurements of bulk- and particle-phase
92 absorption properties have been carried out during a study.



93 Here, we burnt a range of different biomass fuels under different combustion conditions and the
94 resulting aerosol emissions were passed through various in-situ instruments while simultaneously
95 being collected on quartz-fiber filters. The particle phase absorption coefficient was obtained
96 using integrated photoacoustic-nephelometer spectrometers (IPNs) at wavelengths 375, 405 and
97 1047 nm. Organics collected on quartz-fiber filters were extracted in water, acetone, and
98 methanol, and corresponding $b_{\text{abs,bulk}}$ values were calculated. These values were compared with
99 corresponding $b_{\text{abs,OA}}$, and the change in $b_{\text{abs,OA}}/b_{\text{abs,bulk}}$ with varying single scattering albedo
100 (SSA) values and OC/TC ratios was established. SSA was parametrized with the OC/TC ratios
101 with trends similar to those observed by Pokhrel et al., (2016). AÅE from spectroscopic data for
102 bulk and particle phase measurements were compared, and the Mie Theory based correction
103 factor was also verified for a few samples.

104 **2 Methods:**

105 **2.1 Sample generation and collection**

106 Fig. 1 is a schematic diagram of our experimental setup, which consists of a sealed 21 m³
107 stainless-steel combustion chamber housing a fan for mixing and recirculation (Sumlin et al.,
108 2018b). Aerosol samples were generated by burning several types of biomass including pine, fir,
109 grass, sage, and cattle dung (sources are given in the Supplementary Information). During a
110 chamber burn, 10-50 g of a given biomass was placed in a stainless-steel pan and ignited by a
111 butane lighter. The chamber exhaust was kept closed for the duration of a given experiment. The
112 biomass bed was either allowed to burn to completion or it was prematurely extinguished and
113 brought to a smoldering phase by extinguishing the flame beneath a lid. The different
114 combustion conditions were used to generate samples with varying properties: OC/TC ratios



115 ranged from 0.55-1, and SSA values ranged from 0.56-0.98 for wavelengths of 375, 405, and
116 1047 nm.

117 For one set of experiments, the particles were directly sampled from the chamber; in another set,
118 the sampling was done from a hood placed over the burning biomass. A diffusion dryer removed
119 excess water from the sample stream, and the gas-phase organics were removed by a pair of
120 activated parallel-plate semi-volatile organic carbon (SVOC) denuders. The gas-phase organics
121 were stripped to reduce artifacts produced by the adsorption of organic vapors on the quartz
122 filters. The aerosols were finally sent to a 208-liter stainless-steel barrel, from which they were
123 continuously sampled by the three IPNs and a scanning mobility particle sizer (SMPS, TSI, Inc.).
124 The SMPS was not used in all the experiments, however, it gave an estimate of the range over
125 which the size distributions varied and was used to determine the geometric mean of the size
126 distribution. The real-time absorption and scattering coefficients were measured by the IPNs, and
127 samples were simultaneously collected on quartz fiber filters once a steady state signal was
128 achieved. The absorption and scattering coefficients were used to calculate the SSA, which is
129 simply the scattering coefficient divided by the extinction coefficient. Radiative forcing
130 calculations for absorbing OC require good estimates of OC absorption at different SSA values
131 (Lin et al, 2014; Feng et al, 2013; Chakrabarty et al, 2010) underscoring the need for correction
132 factors as a function of SSA. The particles were passed through the filter samplers at a flowrate
133 of 5 lmin^{-1} , with sampling times ranging from 2-15 minutes. Two or more filters were collected
134 for each steady state condition. One of these filters was used to determine the OC and EC
135 fractions of the deposited particles, and the other filters were used for the extraction experiments.

136 **2.2 Analytical Techniques**

137 **2.2.1 Absorption by solvent extracted OC**



138 Quartz filters (Pallflex Tissuquartz) collected during sampling were split into four quarters, and
139 each quarter was extracted using either deionized water, acetone, hexane, or methanol. The
140 absorption by hexane extracts were low and prone to errors, so data for its extracts were not
141 analyzed. The filters were placed in a covered beaker along with 3-5 ml of the solvent for 24
142 hours. The solvent volumes were measured both before and after the extraction and the
143 differences between the two measurements were within 8%. The extracts were then passed
144 through syringe filters with 0.22 μm pores to remove any impurities introduced by the extraction
145 process.

146 The light absorbance of the extracts was measured using a UV-Vis spectrophotometer (Varian
147 Inc., Cary 50) at wavelengths from 300 nm to 800 nm. To compare the absorbance ($A(\lambda)$) of the
148 chromophores in the solution with the absorption coefficient of the particles in the atmosphere,
149 all absorbance values were converted to solution-phase absorption coefficients at given
150 wavelengths ($b_{\text{abs,bulk}}(\lambda)$):

$$151 \quad b_{\text{abs,bulk}}(\lambda) = (A(\lambda) - A(700)) \frac{V_l}{V_a * l} \cdot \ln(10), \quad (1)$$

152 where V_l is the volume of solvent the filter was extracted into, V_a is the volume of air that passed
153 over the given filter area, and l is the optical path length that the beam traveled through the
154 cuvette (1 cm). The absorbance at a given wavelength is normalized to the absorbance at 700 nm
155 to account for any signal drift within the instrument. The resulting absorption coefficient (m^{-1})
156 was multiplied by $\ln(10)$ to convert from base 10 (provided by the UV-Vis spectrophotometer) to
157 natural logarithms.

158 **2.2.2 Absorption by BC and OC in particle phase**



159 To estimate the BC absorption at 375 nm and 405 nm, the absorption data from the IPN operated
160 in the infrared regime at a wavelength of 1047 nm was converted to equivalent particulate
161 absorption at the near UV wavelengths, using a BC absorption Ångström exponent ($A\ddot{A}E_{BC}$)
162 value of 1 (Kirchstetter et al., 2004; Andreae and Gelencsér, 2006). The assumption here being
163 that all the absorption at 1047 nm could be attributed to BC aerosol (Bahadur et al., 2012). The
164 BC light absorption coefficient at shorter wavelengths ($b_{abs,BC}(\lambda_1)$) was calculated by:

$$165 \quad b_{abs,BC}(\lambda_1) = b_{abs,IPN}(1047) \cdot \left(\frac{\lambda_1}{1047}\right)^{-A\ddot{A}E_{BC}}, \quad (2)$$

166 where λ_1 is the wavelength at which the absorption will be calculated and $A\ddot{A}E$ is defined for a
167 pair of wavelengths λ_1 and λ_2 as the exponent in a power law expressing the ratio of the
168 absorption coefficients as follows (Moosmüller et al., 2009):

$$169 \quad A\ddot{A}E(\lambda_1\lambda_2) = \frac{\ln\left[b_{abs}(\lambda_1)/b_{abs}(\lambda_2)\right]}{\ln[\lambda_2/\lambda_1]}, \quad (3)$$

170 $A\ddot{A}E$ is an optical descriptor of the inherent material property. For BC particles, typical values of
171 $A\ddot{A}E \approx 1$, while for OC particles $A\ddot{A}E > 4$ (Moosmüller et al., 2009). The value of $b_{abs,BC}$ at
172 375nm and 405nm was then subtracted from $b_{abs,tot}$ at those wavelengths to calculate $b_{abs,OA}$. The
173 ratio $b_{abs,OA}(\lambda)/b_{abs,bulk}(\lambda)$ was calculated to give the conversion factor for converting from the
174 bulk solvent phase absorption to OA absorption coefficient.

175 The organic and elemental carbon compositions of the filters were measured with a thermal-
176 optical OC/EC analyzer (Sunset Laboratory, Tigard, OR) using the Interagency Monitoring of
177 Protected Visual Environments (IMPROVE)-A analysis method (Chow et al., 2007a). The



178 OC/TC ratios were assumed to be constant for a given steady state IPN reading, which allowed
179 us to relate the absorption data to the OC/TC data.

180 **2.3 Mie Theory Calculations**

181 A commonly used method to correct for differences between the chromophore absorption in
182 solution and aerosol particle absorption is by using Mie Theory (Liu et al., 2013; Washenfelder
183 et al., 2015). The imaginary part (k) of the complex refractive index $m = n + ik$ can be
184 determined from the bulk solution absorption data and converted to an equivalent organic carbon
185 particulate absorption, using Mie Theory along with a range of assumptions.

186 To find k , the mass absorption efficiency (α/ρ) was determined using the absorbance data and the
187 OC mass concentration in the solution:

$$188 \quad \frac{\alpha(\lambda)}{\rho} = \frac{b_{abs,bulk}(\lambda)}{M}, \quad (4)$$

189 where $b_{abs,bulk}(\lambda)$ is the bulk absorption coefficient determined in Eq. (1), and M is the mass
190 concentration of OC in the solution. In the given study, the mass concentration was measured for
191 some of the water extracts using a total organic carbon (TOC) analyzer (Shimadzu, TOC-L). The
192 corresponding water-soluble organic carbon (WSOC) was then used to estimate α/ρ of the bulk
193 solution. The calculated α/ρ was then used to determine k for the WSOC by:

$$194 \quad k(\lambda) = \frac{\rho \cdot \lambda \cdot \left(\frac{\alpha(\lambda)}{\rho}\right)}{4\pi}, \quad (5)$$

195 where λ is the light wavelength at which k needs to be calculated, and ρ is the density of the
196 dissolved organic compounds. A ρ value of 1.6 (Alexander et al., 2008) was used to calculate the
197 k values, and also used in all subsequent calculations using density. A Mie based inversion



198 algorithm was used to extract the real part of the refractive index (n) using data from the SMPS
199 and IPN (Sumlin et al., 2018a). A sensitivity analysis was performed by varying the n value from
200 1.4 to 2, and the change in Mie calculated absorption was within 18%. The size distribution for
201 the WSOC was estimated assuming the same geometric mean and standard deviation as that of
202 the original aerosol, but with number concentrations calculated based on the extracted mass.
203 Calculations for the number concentration are provided in the SI.

204 After the size distribution and complex refractive index were determined, they were used to
205 calculate the absorption coefficient based on Mie Theory, which was then compared to $b_{\text{abs,bulk}}$ to
206 determine Mie based correction factors for converting from bulk to particle phase absorption.

207 3 Results and discussion

208 3.1 Correction factor as a function of Single Scattering Albedo

209 Fig. 2 shows the trends in $b_{\text{abs,OA}}(\lambda)/b_{\text{abs,bulk}}(\lambda)$ with varying SSA. The error bars account for
210 uncertainties in IPN measurements, UV-Vis spectrophotometer measurements, flowrates for
211 filter sampling, and extract volume measurements. The SSA values of fractal BC aggregates are
212 between 0.3-0.4 for light wavelengths of 400 nm (Cheng et al., 2013), and due to this particularly
213 low SSA of BC compared to OC, an increase in the BC content of the aerosol composition is
214 expected with decreasing SSA. This relation is explored further in Section 3.2 and 3.3. Fig 2.
215 also indicates that the light absorbed by methanol and acetone extracts were almost identical and
216 would imply that the amount and type of OC extracted by the two solvents were similar, as seen
217 in other studies as well (Chen and Bond, 2010; Wang et al., 2014). The differences between the
218 conversion factors for water and methanol are discussed further in Section 3.3. The differences



219 between the mean values of $b_{\text{abs,OA}}(\lambda)/b_{\text{abs,bulk}}(\lambda)$ at 375 and 405 nm were less than or close to the
220 errors associated with them, hence any trends with wavelength were not explored.

221 The value of $b_{\text{abs,OA}}(\lambda)/b_{\text{abs,bulk}}(\lambda)$ approached a constant in the measured range of data. A power
222 law ($y = k_0 + k_1 x^{k_2}$) was used to fit the points in Fig. 2, and the corresponding fit parameters,
223 along with root mean square error (RMSE) values, are listed in Table 1. The value of BC ÅÅE
224 has a significant effect on the conversion factor for particles with SSA smaller than 0.7 at 375
225 nm and smaller than 0.825 at 405 nm. A sensitivity analysis was performed to observe the
226 change in $b_{\text{abs,OA}}(\lambda)/b_{\text{abs,bulk}}(\lambda)$ due to a change in BC ÅÅE from 0.85 to 1.1, which are typical
227 values observed in other studies (Lack et al., 2008; Bergstrom et al., 2007; Lan et al., 2013). For
228 the correction factors at 375 nm, on varying the ÅÅE we calculated a change of less than 14%
229 for aerosols with SSA greater than 0.7, whereas a change as high as 200% was seen for particles
230 with SSA less than 0.7. Similar changes in the correction factor were observed for SSA less than
231 0.825 at 405 nm. These fluctuations indicate that the current method is best suited for particles
232 with relatively higher SSA values. The perforated lines in Fig. 2 separate points with high ÅÅE
233 based fluctuations (at lower SSA values) from those that have relatively small variation with
234 changes in the BC ÅÅE.

235 3.2 SSA parametrized with OC/TC

236 Fig. 3 shows the variation in SSA with change in the OC/TC ratio of the aerosol. The OC/TC
237 ratio was determined using the IMPROVE-A TOR protocol with the thermal optical EC/OC
238 analyzer at Sunset laboratories. The resulting fits for the plots in Fig. 3 are linear with correlation
239 coefficients of 0.9 and 0.95 for wavelengths 375 nm and 405 nm respectively. A linear relation
240 between the SSA and the EC/TC ratio (which is simply the OC/TC ratio subtracted from 1) was
241 also observed by Pokhrel et al. (2016). However, when the data from that study were converted



242 to OC/TC values for comparison, it was noted that the slopes and intercepts of the resulting fits
243 were different from those observed in this study. Table 2 has a list of the slope and intercept of
244 fits for comparable wavelengths in both studies, along with the correlation coefficient.

245 In Fig. 3, the points corresponding to high OC/TC ratios are associated with SSA values that are
246 close to 1, because pure OC aerosols are predominantly light-scattering. However, it is also
247 noted that the aerosols with OC/TC of 1 have SSA values less than 1 at the measured
248 wavelengths and the fit values indicate that the absorption has large spectral variation. The fit
249 yielded SSA values of 0.86 and 0.95 at 375 and 405 nm respectively indicating that the aerosols
250 can be categorized as brown carbon (Chakrabarty et al. 2010). Thus, it is useful to note that the
251 fit parameters in Table 2 capture the effects of the brown carbon, because the SSA values for
252 pure OC are below 1 at both wavelengths.

253 A useful conclusion from Fig. 3 is that the OC/TC ratio can predict the SSA of the aerosol even
254 though information on size distributions, burn conditions, and fuel type are not used. This
255 correlation, however, has high variations at OC/TC ratios very close to 1, where fuel type and
256 burn conditions do dictate the composition and absorption properties (Chen and Bond, 2010;
257 Budisulistiorini et al., 2017) of organics released and hence a larger range of SSA values exist
258 for the given OC/TC value. Further studies need to be conducted using more fuels with a variety
259 of distinct size distributions and burn conditions to determine the validity and exact parameters
260 for the fit.

261 3.3 Correction factor as a function of OC/TC ratio

262 Fig. 4 depicts the variation in $b_{\text{abs,OA}}(\lambda)/b_{\text{abs,bulk}}(\lambda)$ with different OC/TC ratios. Because the
263 OC/TC ratio and the SSA are well correlated, we expect to see a similar trend for Fig. 4 as in



264 Fig. 2. The correction factor for the three solvents approaches a constant as the OC/TC ratio
265 decreases. However, it is important to note that the magnitude of the conversion factor could
266 change due to uncertainties in the value of the A_{λ}^{BC} used for the BC. A power law like the one in
267 Fig. 2 was fit to the data in Fig. 4. The fit parameters for the different solvents at the two
268 wavelengths, along with the RMSE value for each fit, are presented in Table 3. The errors due to
269 uncertainty in BC A_{λ}^{BC} are more pronounced at higher EC fractions, leading to high fluctuations
270 in correction factors at OC/TC ratios below 0.75. Thus, the correction factors determined by this
271 method are valid primarily for aerosols with EC fraction < 0.25 .

272 We also fit an exponential curve to $b_{\text{abs,OA}}(\lambda)/b_{\text{abs,bulk}}(\lambda)$ as a function of EC/OC values ($y = k_0 +$
273 $k_1 * (1 - \exp(-(EC/OC)/k_2))$). The fits had comparable RMSE values, but because strong
274 relationships between the SSA and OC/TC ratios were previously established, the plots described
275 here continue along those lines. The curves plotted with EC/OC, along with the parameters for
276 the fits, are provided in the SI.

277 The differences in the magnitudes of the correction factors between acetone/methanol extracts
278 and water extracts increase as the EC composition of the aerosols increases. A change in the
279 absorption properties of OC at different EC/OC ratios was observed by Saleh et al. (2014) which
280 could explain the increasing difference in the magnitude of $b_{\text{abs,OA}}/b_{\text{abs,bulk}}$ between water and
281 methanol/acetone extracts. The increasing difference in absorbance between water and methanol
282 with increasing EC content would indicate a decrease in the total absorbing WSOC compared to
283 total absorbing OC. An increase in the amount of water insoluble organic carbon with increasing
284 EC content of the aerosol was also observed by Zhang et al. (2013), which strengthens our
285 hypothesis for a steeper increase in $b_{\text{abs,OA}}(\lambda)/b_{\text{abs,bulk}}(\lambda)$ of water extracts as compared to
286 methanol with increasing BC mass fraction of the aerosol.



287 3.4 Variations in AÅE with solvents and OC/TC ratios

288 The AÅE values, for organics extracted in different solvents and those obtained from $b_{\text{abs,OA}}$ are
289 compared in Table 4. The AÅE values were calculated between $\lambda = 375$ and 405 nm from
290 spectrophotometer data. Consistent with previous studies (Chen and Bond, 2010; Zhang et al.,
291 2013; Liu et al., 2013), the AÅE values of the water extracts were larger than the AÅE of the
292 acetone and methanol extracts. The lower AÅE values for the methanol and acetone extracts
293 compared to water would indicate that the former two solvents are better at extracting the higher
294 molecular weight compounds than water, and they can thus absorb more light at longer
295 wavelengths, leading to smaller AÅE.

296 The AÅE values calculated for OA ranged from 4.4 to 14.61 and overall the AÅE for OA
297 measurements decreased with increasing EC mass fraction in the aerosol. The AÅE calculated
298 for OC in both the particle- and bulk-phase decreased at lower OC/TC ratios indicating a drastic
299 decrease in the spectral dependence of OC co-emitted with BC. Table 4 shows that AÅE values
300 for OA were greater than or close to their bulk counterparts at OC/TC ratios close to 1. These
301 bulk measurements of AÅE suggest that they deviate significantly from the spectral dependence
302 of OC in the particle phase, and future studies and models should not use AÅE data from bulk
303 measurements to be representative of the particle phase.

304 3.4 Mie Calculations

305 The absorption coefficient determined from the bulk absorbance using Eq. (1) was compared to
306 the absorption coefficient calculated using Mie theory for three samples of smoldering sage. The
307 EC/OC analysis (IMPROVE-A protocol) determined that these samples consisted purely of OC,



308 and because the SMPS measurements and TOC analysis were not performed on all samples, the
309 three samples of sage were considered optimum for the Mie calculations.

310 The Mie based correction factors for bulk to particulate absorption for the three samples are
311 presented in Table 5. The calculated correction factors at 375 nm and 405 nm are close to 2 as
312 observed in previous studies (Liu et al., 2013; Washenfelder et al., 2015). The values for the
313 correction factor vary from 1.99 to 2.05 at 375 nm and 2.15 to 2.29 at 405 nm. However, it is
314 important to note that the correction factor of 2 cannot be used for all conditions, as observations
315 from Fig. 2 and Fig. 4 indicate that the Mie theory based correction factor might severely
316 underestimate the absorption by the total absorbing OA particles in accumulation mode.

317 **4 Conclusions**

318 Under controlled laboratory conditions, we determined the factor required for converting the
319 absorption coefficient of bulk solvent extracts to particle phase absorption coefficients for
320 organic aerosol emissions from biomass combustion. We combusted a range of different
321 wildland fuels under different combustion conditions, generating a span of different SSA and
322 OC/TC values. The SSA values ranged from 0.55 to 0.87 at 375 nm, and from 0.69 to 0.95 at
323 405 nm, the OC/TC values ranged from 0.55 to 1. We observed that the conversion factor tends
324 towards a constant with increasing EC content for the range of SSA and OC/TC analyzed, and
325 these factors were parametrized with the SSA and OC/TC of the total aerosol for three solvents
326 (water, methanol, and acetone). We also demonstrated that the SSA and OC/TC ratios can be
327 well parametrized with a linear fit that captures the effects of brown carbon aerosol. We analyzed
328 the validity of the conventionally used correction factor of 2 for water extracts of aerosol
329 particles in accumulation mode and noted that, while the factor is reproducible, its use can
330 significantly underestimate OA absorption. We recommend using $b_{\text{abs,OA}}/b_{\text{abs,bulk}}$ values between



331 2 and 11 for water extracts and values between 1 and 4 for methanol extracts based on OC/TC
332 ratios for EC mass fractions less than 0.25.

333 For future studies, a better technique to quantify BC absorption at lower wavelengths, such as a
334 thermodenuder to strip off all OC, or a single particle soot photometer along with core-shell Mie
335 calculations to determine BC absorption can be used to decrease uncertainties from BC $\Delta\Delta E$.
336 Here, BC is assumed to be externally mixed with OC; correction factors accounting for BC
337 absorption enhancement can be explored to improve estimates of $b_{\text{abs,OA}}/b_{\text{abs,bulk}}$. Zhang et al.
338 (2013) observed lower $\Delta\Delta E$ for WSOC from a particle into liquid sampler (PILS) than for
339 methanol extracts. The hypothesis was that the highly dilute environment in PILS increased
340 dissolution of organics in water. The validity and reason for lower $\Delta\Delta E$ from PILS data can be
341 explored in future studies along with finding corresponding differences in $b_{\text{abs,OA}}/b_{\text{abs,bulk}}$.

342 **Author Contributions**

343 RKC conceived of this study and designed the experiments. SB and WMH collected the fuels for
344 the experiments and performed EC/OC analysis on the sampled filters. NS and AP carried out
345 the experiments and analysed the data. NS prepared the manuscript with input from all co-
346 authors.

347 **Acknowledgements**

348 This work was partially supported by the National Science Foundation under Grant No.
349 AGS1455215, NASA ROSES under Grant No. NNX15AI66G.



350 References

- 351 Alexander, D. T., Crozier, P. A., and Anderson, J. R.: Brown carbon spheres in East Asian
352 outflow and their optical properties, *Science*, 321, 833-836, 2008.
- 353 Andreae, M., and Gelencsér, A.: Black carbon or brown carbon? The nature of light-absorbing
354 carbonaceous aerosols, *Atmospheric Chemistry and Physics*, 6, 3131-3148, 2006.
- 355 Arnott, W., Moosmüller, H., Sheridan, P., Ogren, J., Raspert, R., Slaton, W., Hand, J.,
356 Kreidenweis, S., and Collett Jr, J.: Photoacoustic and ~~biased~~ ambient aerosol light
357 absorption measurements: Instrument comparisons and the role of relative humidity, *Journal of*
358 *Geophysical Research: Atmospheres*, 108, AAC 15-11-AAC 15-11, 2003.
- 359 Arnott, W. P., Hamasha, K., Moosmüller, H., Sheridan, P. J., and Ogren, J. A.: Towards aerosol
360 light-absorption measurements with a 7-wavelength aethalometer: Evaluation with a
361 photoacoustic instrument and 3-wavelength nephelometer, *Aerosol Science and Technology*, 39,
362 17-29, 2005.
- 363 Arola, A., Schuster, G., Myhre, G., Kazadzis, S., Dey, S., and Tripathi, S.: Inferring absorbing
364 organic carbon content from AERONET data, *Atmospheric Chemistry and Physics*, 11, 215-225,
365 2011.
- 366 Bahadur, R., Praveen, P. S., Xu, Y., and Ramanathan, V.: Solar absorption by elemental and
367 brown carbon determined from spectral observations, *Proceedings of the National Academy of*
368 *Sciences*, 201205910, 2012.
- 369 Bergstrom, R. W., Pilewskie, P., Russell, P. B., Redemann, J., Bond, T. C., Quinn, P. K., and
370 Sierau, B.: Spectral absorption properties of atmospheric aerosols, *Atmospheric Chemistry and*
371 *Physics*, 7, 5937-5943, 2007.
- 372 Bond, T. C., Doherty, S. J., Fahey, D., Forster, P., Berntsen, T., DeAngelo, B., Flanner, M.,
373 Ghan, S., Kärcher, B., and Koch, D.: Bounding the role of black carbon in the climate system: A
374 scientific assessment, *Journal of Geophysical Research: Atmospheres*, 118, 5380-5552, 2013.
- 375 Budisulistiorini, S. H., Riva, M., Williams, M., Chen, J., Itoh, M., Surratt, J. D., and Kuwata, M.:
376 Light-absorbing brown carbon aerosol constituents from combustion of Indonesian peat and
377 biomass, *Environmental Science & Technology*, 51, 4415-4423, 2017.
- 378 Chakrabarty, R., Moosmüller, H., Chen, L.-W., Lewis, K., Arnott, W., Mazzoleni, C., Dubey,
379 M., Wold, C., Hao, W., and Kreidenweis, S.: Brown carbon in tar balls from smoldering biomass
380 combustion, *Atmospheric Chemistry and Physics*, 10, 6363-6370, 2010.
- 381 Chen, Y., and Bond, T.: Light absorption by organic carbon from wood combustion,
382 *Atmospheric Chemistry and Physics*, 10, 1773-1787, 2010.



- 383 Chen, Y., Wang, H., Singh, B., Ma, P. L., Rasch, P. J., and Bond, T. C.: Investigating the linear
384 dependence of direct and indirect radiative forcing on emission of carbonaceous aerosols in a
385 global climate model, *Journal of Geophysical Research: Atmospheres*, 123, 1657-1672, 2018.
- 386 Cheng, T., Gu, X., Wu, Y., Chen, H., and Yu, T.: The optical properties of absorbing aerosols
387 with fractal soot aggregates: Implications for aerosol remote sensing, *Journal of Quantitative
388 Spectroscopy and Radiative Transfer*, 125, 93-104, 2013.
- 389 Chow, J. C., Watson, J. G., Chen, L.-W. A., Chang, M. O., Robinson, N. F., Trimble, D., and
390 Kohl, S.: The IMPROVE_A temperature protocol for thermal/optical carbon analysis:
391 maintaining consistency with a long-term database, *Journal of the Air & Waste Management
392 Association*, 57, 1014-1023, 2007a.
- 393 Chow, J. C., Yu, J. Z., Watson, J. G., Hang Ho, S. S., Bohannan, T. L., Hays, M. D., and Fung,
394 K. K.: The application of thermal methods for determining chemical composition of
395 carbonaceous aerosols: A review, *Journal of Environmental Science and Health Part A*, 42,
396 1521-1541, 2007b.
- 397 Chung, C. E., Ramanathan, V., and Decremer, D.: Observationally constrained estimates of
398 carbonaceous aerosol radiative forcing, *Proceedings of the National Academy of Sciences*, 109,
399 11624-11629, 2012.
- 400 Feng, Y., Ramanathan, V., and Kotamarthi, V.: Brown carbon: a significant atmospheric
401 absorber of solar radiation?, *Atmospheric Chemistry and Physics*, 13, 8607-8621, 2013.
- 402 Husain, L., Dutkiewicz, V. A., Khan, A., and Ghauri, B. M.: Characterization of carbonaceous
403 aerosols in urban air, *Atmospheric Environment*, 41, 6872-6883, 2007.
- 404 IPCC: *Climate Change: The Physical Science Basis, Contribution of Working Group I to the UN
405 IPCC's 5th Assessment Report*, Cambridge University Press, New York, USA, 2013.
- 406 Kirchstetter, T. W., Novakov, T., and Hobbs, P. V.: Evidence that the spectral dependence of
407 light absorption by aerosols is affected by organic carbon, *Journal of Geophysical Research:
408 Atmospheres*, 109, 2004.
- 409 Kirchstetter, T., and Thatcher, T.: Contribution of organic carbon to wood smoke particulate
410 matter absorption of solar radiation, *Atmospheric Chemistry and Physics*, 12, 6067-6072, 2012.
- 411 Lack, D. A., Lovejoy, E. R., Baynard, T., Pettersson, A., and Ravishankara, A.: Aerosol
412 absorption measurement using photoacoustic spectroscopy: Sensitivity, calibration, and
413 uncertainty developments, *Aerosol Science and Technology*, 40, 697-708, 2006.
- 414 Lack, D. A., Cappa, C. D., Covert, D. S., Baynard, T., Massoli, P., Sierau, B., Bates, T. S.,
415 Quinn, P. K., Lovejoy, E. R., and Ravishankara, A.: Bias in filter-based aerosol light absorption
416 measurements due to organic aerosol loading: Evidence from ambient measurements, *Aerosol
417 Science and Technology*, 42, 1033-1041, 2008.



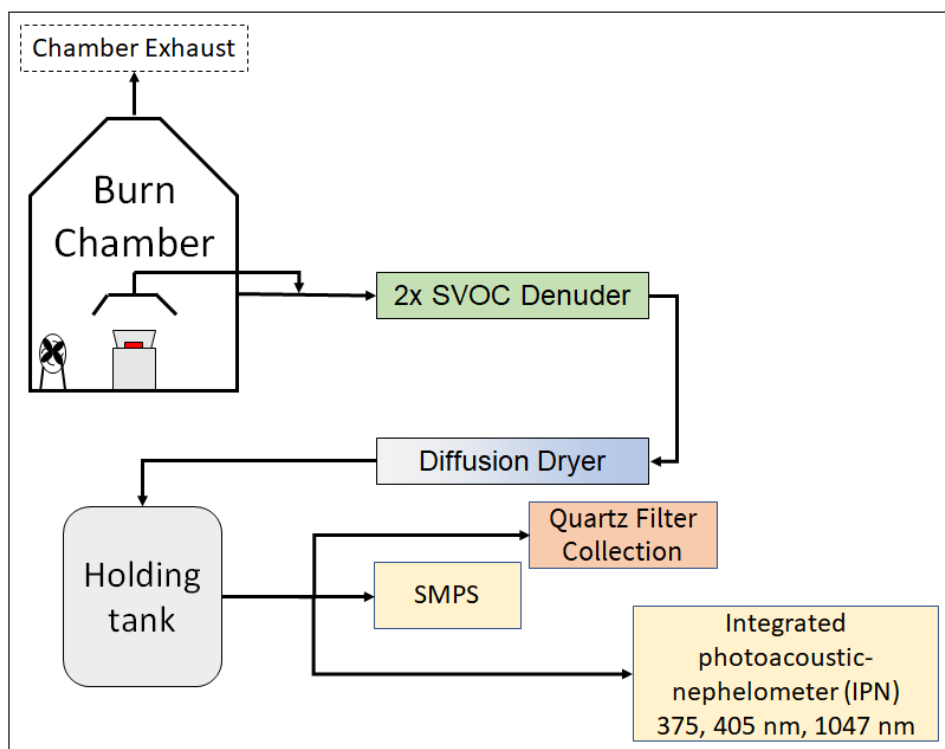
- 418 Lan, Z.-J., Huang, X.-F., Yu, K.-Y., Sun, T.-L., Zeng, L.-W., and Hu, M.: Light absorption of
419 black carbon aerosol and its enhancement by mixing state in an urban atmosphere in South
420 China, *Atmospheric environment*, 69, 118-123, 2013.
- 421 Lin, G., Penner, J. E., Flanner, M. G., Sillman, S., Xu, L., and Zhou, C.: Radiative forcing of
422 organic aerosol in the atmosphere and on snow: Effects of SOA and brown carbon, *Journal of*
423 *Geophysical Research: Atmospheres*, 119, 7453-7476, 2014.
- 424 Liu, J., Bergin, M., Guo, H., King, L., Kotra, N., Edgerton, E., and Weber, R.: Size-resolved
425 measurements of brown carbon in water and methanol extracts and estimates of their
426 contribution to ambient fine-particle light absorption, *Atmospheric Chemistry and Physics*, 13,
427 12389-12404, 2013.
- 428 Liu, J., Scheuer, E., Dibb, J., Ziemba, L. D., Thornhill, K. L., Anderson, B. E., Wisthaler, A.,
429 Mikoviny, T., Devi, J. J., and Bergin, M.: Brown carbon in the continental troposphere,
430 *Geophysical Research Letters*, 41, 2191-2195, 2014.
- 431 Liu, X., Zhang, Y., Huey, L., Yokelson, R., Wang, Y., Jimenez, J., Campuzano, P.,
432 Beyersdorf, A., Blake, D., and Choi, Y.: Agricultural fires in the southeastern US during
433 SEAC4RS: Emissions of trace gases and particles and evolution of ozone, reactive nitrogen, and
434 organic aerosol, *Journal of Geophysical Research: Atmospheres*, 121, 7383-7414, 2016.
- 435 Mo, Y., Li, J., Liu, J., Zhong, G., Cheng, Z., Tian, C., Chen, Y., and Zhang, G.: The influence of
436 solvent and pH on determination of the light absorption properties of water-soluble brown
437 carbon, *Atmospheric environment*, 161, 90-98, 2017.
- 438 Moosmüller, H., Chakrabarty, R., and Arnott, W.: Aerosol light absorption and its measurement:
439 A review, *Journal of Quantitative Spectroscopy and Radiative Transfer*, 110, 844-878, 2009.
- 440 Moosmüller, H., Chakrabarty, R., Ehlers, K., and Arnott, W.: Absorption Ångström coefficient,
441 brown carbon, and aerosols: basic concepts, bulk matter, and spherical particles, *Atmospheric*
442 *Chemistry and Physics*, 11, 1217-1225, 2011.
- 443 Pokhrel, R. P., Wagner, N. L., Langridge, J. M., Lack, D. A., Jayarathne, T., Stone, E. A.,
444 Stockwell, C. E., Yokelson, R. J., and Murphy, S. M.: Parameterization of single-scattering
445 albedo (SSA) and absorption Ångström exponent (AAE) with EC/OC for aerosol emissions from
446 biomass burning, *Atmospheric Chemistry and Physics*, 16, 9549-9561, 2016.
- 447 Ramanathan, V., and Carmichael, G.: Global and regional climate changes due to black carbon,
448 *Nature geoscience*, 1, 221, 2008.
- 449 Saleh, R., Robinson, E. S., Tkacik, D. S., Ahern, A. T., Liu, S., Aiken, A. C., Sullivan, R. C.,
450 Presto, A. A., Dubey, M. K., and Yokelson, R. J.: Brownness of organics in aerosols from
451 biomass burning linked to their black carbon content, *Nature Geoscience*, 7, 647, 2014.



- 452 Sumlin, B. J., Heinson, W. R., and Chakrabarty, R. K.: Retrieving the aerosol complex refractive
453 index using PyMieScatt: A Mie computational package with visualization capabilities, *Journal of*
454 *Quantitative Spectroscopy and Radiative Transfer*, 205, 127-134, 2018a.
- 455 Sumlin, B. J., Heinson, Y. W., Shetty, N., Pandey, A., Pattison, R. S., Baker, S., Hao, W. M., and
456 Chakrabarty, R. K.: UV-Vis-IR spectral complex refractive indices and optical properties of
457 brown carbon aerosol from biomass burning, *Journal of Quantitative Spectroscopy and Radiative*
458 *Transfer*, 206, 392-398, 2018b.
- 459 Sun, H., Biedermann, L., and Bond, T. C.: Color of brown carbon: A model for ultraviolet and
460 visible light absorption by organic carbon aerosol, *Geophysical Research Letters*, 34, 2007.
- 461 Wang, X., Heald, C., Ridley, D., Schwarz, J., Spackman, J., Perring, A., Coe, H., Liu, D., and
462 Clarke, A.: Exploiting simultaneous observational constraints on mass and absorption to estimate
463 the global direct radiative forcing of black carbon and brown carbon, *Atmospheric Chemistry*
464 *and Physics*, 14, 10989-11010, 2014.
- 465 Washenfelder, R., Attwood, A., Brock, C., Guo, H., Xu, L., Weber, R., Ng, N., Allen, H., Ayres,
466 B., and Baumann, K.: Biomass burning dominates brown carbon absorption in the rural
467 southeastern United States, *Geophysical Research Letters*, 42, 653-664, 2015.
- 468 Yang, M., Howell, S., Zhuang, J., and Huebert, B.: Attribution of aerosol light absorption to
469 black carbon, brown carbon, and dust in China—interpretations of atmospheric measurements
470 during EAST-AIRE, *Atmospheric Chemistry and Physics*, 9, 2035-2050, 2009.
- 471 Zhang, X., Wang, Y., Zhang, X., Guo, W., and Gong, S.: Carbonaceous aerosol composition
472 over various regions of China during 2006, *Journal of Geophysical Research: Atmospheres*, 113,
473 2008.
- 474 Zhang, X., Lin, Y.-H., Surratt, J. D., and Weber, R. J.: Sources, composition and absorption
475 Ångström exponent of light-absorbing organic components in aerosol extracts from the Los
476 Angeles Basin, *Environmental science & technology*, 47, 3685-3693, 2013.



477 **Figures and Tables:**



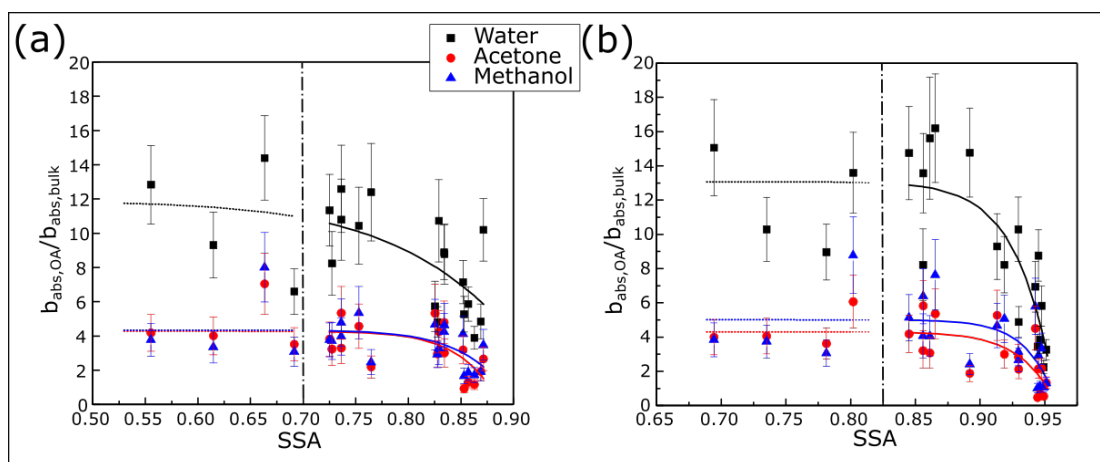
478

479 **Fig. 1:** A schematic representing the experimental setup. The aerosol emissions were either
 480 sampled directly from the chamber wall or through a hood placed directly above the combusting
 481 biomass.

482



483



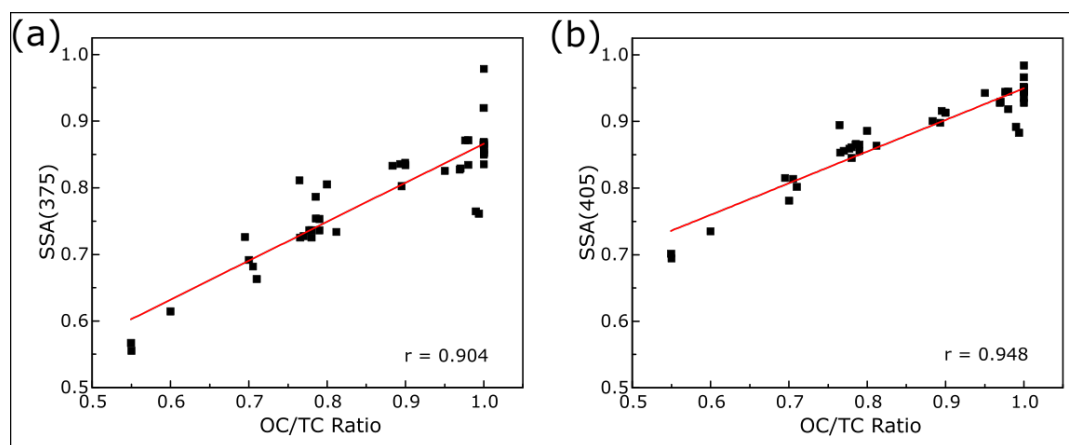
484

485 **Fig. 2:** Variation in $b_{\text{abs,OA}}/b_{\text{abs,bulk}}$ with change in the SSA at (a) 375 nm and (b) 405 nm. The
486 error bars represent one standard deviation from the mean. The perforated lines separate points at
487 lower SSA, which have high fluctuations with changing BC ÅÅE, from the data at high SSA.

488



489



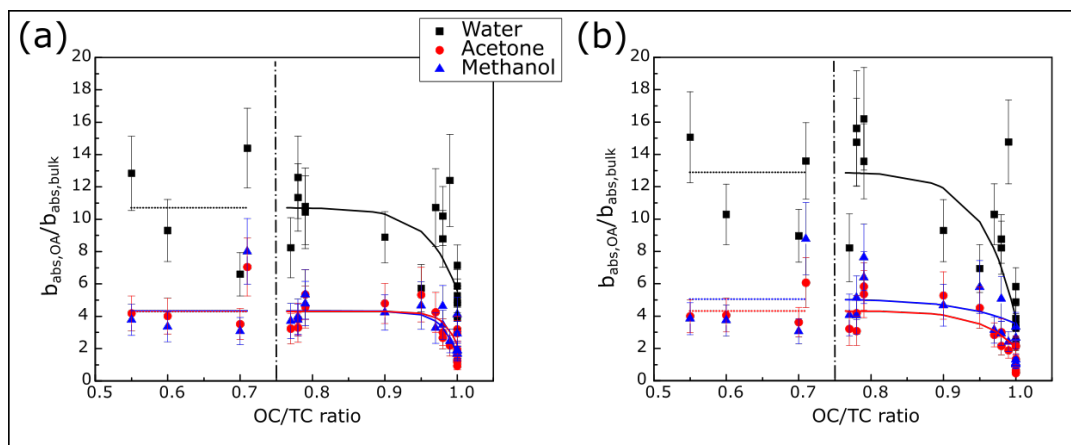
490

491 **Fig. 3:** SSA at (a) 375 nm and (b) 405 nm as a function of the OC/TC ratio. The solid red lines
492 are linear fits to the data. Pearson's r values are given at the bottom right in each plot, and the
493 parameters for the fits are listed in Table 2.

494



495



496

497 **Fig. 4:** The values of $b_{\text{abs,OA}}/b_{\text{abs,bulk}}$ plotted with the OC/TC ratio, instead of the SSA, as in Fig.

498 2.



499 **Table 1:** Fit coefficients for $b_{\text{abs},\text{OA}}(\lambda)/b_{\text{abs},\text{bulk}}(\lambda)$ as a function of SSA ($y = k_0 + k_1 (\text{SSA})^{k_2}$)
 500 for tested solvents and the fuels analyzed in this study along with the RMSE value for each fit.

	Wavelength (nm)	Solvent	Fit Parameters			RMSE
			k_0	k_1	k_2	
$\frac{b_{\text{abs},\text{OA}}}{b_{\text{abs},\text{bulk}}}$	375	Water	11.84	-19.09	8.43	2.27
		Acetone	4.28	-50.8	24.52	1.25
		Methanol	4.34	-41.01	21.41	2.15
	405	Water	13.07	-52.3	33.34	2.46
		Acetone	4.32	-18.69	35.19	1.17
		Methanol	5.00	-29.47	42.97	1.63

501



502 **Table 2.** Fit coefficients along with 95% confidence intervals in brackets for plots of SSA v/s
503 OC/TC ratios ($y = m \text{ (OC/TC)} + c$) for the different biomass fuels used in this study, and
504 parameters for fits from Pokhrel et al. (2016) for 405 nm, along with the Pearson's r value for
505 each linear fit.

	Wavelength (nm)	m	c	r
This study	375	0.585 (± 0.040)	0.281 (± 0.036)	0.904
	405	0.475 (± 0.023)	0.475 (± 0.021)	0.948
Pokhrel	405	0.874 (± 0.049)	0.036 (± 0.042)	0.971

506



507 **Table 3:** Fit parameters for ratios of the absorption coefficient of organics in the particle phase to
 508 the absorption coefficient of the bulk phase, as a function of the OC/TC ratio ($y = k_0 +$
 509 $k_1(OC/TC)^{k_2}$) for the fuels analyzed in this study, along with the RMSE value for each fit.

	Wavelength (nm)	Solvent	Fit Parameters			
			k_0	k_1	k_2	RMSE
$\frac{b_{abs,OA}}{b_{abs,bulk}}$	375	Water	10.71	-5.13	24.32	2.11
		Acetone	4.26	-2.6	48.11	1.03
		Methanol	4.26	-2	53.6	2.14
	405	Water	12.89	-9.02	21.06	2.34
		Acetone	4.29	-3.28	31.5	0.97
		Methanol	4.96	-3.39	40.85	1.51

510



511 **Table 4:** The $\Delta A E$ of OA from various fuels extracted in water, acetone, and methanol, along
 512 with the $\Delta A E$ calculated for $b_{abs,OA}$.

Fuel	OC/TC ratio	$\Delta A E_{375-405}$			
		OA	Water	Acetone	Methanol
Dung	1	13.13	8.00	5.29	5.21
	1	14.44	8.95	5.85	7.75
	1	14.61	7.48	4.62	4.5
	1	14.08	8.55	5.25	6.8
Sage	1	13.53	10.87	8.62	8.8
	1	10.54	10.71	6.29	7.3
	0.97	10.41	9.88	5.2	5.8
	0.79	6.57	12.28	8.62	9.17
	0.79	7.6	10.58	8.73	8.3
	0.71	8.23	7.5	6.29	6.4
	0.55	4.4	6.48	3.84	3.6
Grass	0.99	9.8	12.08	7.84	7.5
	0.78	9.35	10.15	8.47	9.6
	0.78	6.3	9.7	7.49	7.3
	0.77	8.22	8.21	8.12	8.4
Pine	0.98	11.35	9.36	8.55	8.1
	0.98	8.43	9.59	8.36	8.6
	0.95	13.97	16.44	11.81	12.8
	0.9	7.51	9.09	8.75	8.7
	0.7	5.96	9.93	6.33	5.83
	0.6	5.04	6.36	5.2	5.4

513



514 **Table 5:** Correction factors for bulk solution absorption to particle phase absorption, based on
515 Mie Theory calculations.

Fuel	Geometric mean (in nm)	Geometric standard deviation	Mie based Correction Factor	
			375 nm	405 nm
Sage	397	1.3	2.04 ± 0.38	2.27 ± 0.41
	271	1.32	2.05 ± 0.38	2.29 ± 0.41
	159	1.59	1.99 ± 0.36	2.15 ± 0.39

516