| 1 | Measuring Light Absorption by Freshly Emitted Organic Aerosols: Optical Artifacts in |
|--------|---|
| 2 | Traditional Solvent Extraction-Based Methods |
| 3 | Nishit J Shetty ¹ , Apoorva Pandey ¹ , Stephen Baker ² , Wei Min Hao ² , Rajan K. Chakrabarty ^{1,3} |
| 4 5 | ¹ Center for Aerosol Science and Engineering, Department of Energy, Environmental and Chemical Engineering, Washington University in St. Louis, St. Louis, MO 63130, USA |
| 6 7 | ² USDA Forest Service, Rocky Mountain Research Station, Fire Sciences Laboratory, Missoula, Montana, USA |
| 8 9 | ³ McDonnell Center for the Space Sciences, Washington University in St. Louis, St. Louis, MO 63130, USA |
| 10 | Correspondence to: Rajan K. Chakrabarty (chakrabarty@wustl.edu) |
| 11 | Abstract |
| 12 | Recent studies have shown that organic aerosol (OA) could have a non-trivial role in atmospheric |
| 13 | light absorption at shorter visible wavelengths. Good estimates of OA absorption are therefore |
| 14 | necessary to accurately calculate radiative forcing due to these aerosols in climate models. One of |
| 15 | the common techniques used to measure OA light absorption is the solvent extraction technique |
| 16 | from filter samples which involves the use of a spectrophotometer to measure bulk absorbance by |
| 17 | the solvent-soluble organic fraction of particulate matter. Measured solvent phase absorbance is |
| 18 | subsequently converted to particle-phase absorption coefficient using scaling factors. The |
| 19 | conventional view is to apply a correction factor of 2 to absorption coefficients obtained from |

biases due to incomplete extraction of OC by solvents and size-dependent absorption properties of
OA. The range for these biases along with their potential dependence on burn conditions is an

solvent-extracted OA based on Mie calculations. The appropriate scaling factors are a function of

23 unexplored area of research.

Here, we performed a comprehensive laboratory study involving three solvents (water, methanol, 24 and acetone) to investigate the bias in absorption coefficients obtained from solvent extraction-25 26 based photometry techniques as compared to in-situ particle phase absorption for freshly emitted OA from biomass burning. We correlated the bias with OC/TC mass ratio and single scattering 27 albedo (SSA) and observed that the conventionally used correction factor of 2 for water and 28 29 methanol-extracted OA might not be extensible to all systems and suggest caution while using such correction factors to estimate particle-phase OA absorption coefficients. Furthermore, a linear 30 31 correlation between SSA and OC/TC ratio was also established. Finally, from the spectroscopic data, we analyzed the differences in Absorption Ångström Exponents (AÅE) obtained from 32 solution- and particulate-phase measurements. We noted that AAE from solvent phase 33 34 measurements could deviate significantly from their OA counterparts.

35 **1 Introduction**

Carbonaceous aerosols constitute a major short-lived climate pollutant, and even though they have 36 37 been studied extensively in recent years, estimates of their contribution to shortwave radiative forcing remains highly uncertain (IPCC, 2013). Based on their thermal-refractory properties, 38 carbonaceous aerosols are categorized as elemental carbon (EC) or organic carbon (OC) (Chow et 39 40 al., 2007b; Bond et al., 2013), and the sum of OC and EC is referred to as total carbon (TC). When defined optically, the refractory EC component is approximately referred to as black carbon (BC) 41 42 (Chow et al., 2007b; Bond et al., 2013); BC aerosol constitute the strongest of the light absorbing aerosol components in the atmosphere (Ramanathan and Carmichael, 2008; Andreae and 43 Gelencsér, 2006; IPCC, 2013). While BC absorbs strongly in the visible spectrum, the contribution 44 of OC towards absorption has largely been neglected, even though many studies have 45 demonstrated significant OC absorption at lower visible wavelengths (Yang et al., 2009; Chen and 46

Bond, 2010; Chakrabarty et al., 2010; Kirchstetter 2012). The atmospheric mass of OC can be 312 times larger than that of BC (Husain et al., 2007; Zhang et al., 2008) which warrants its inclusion
as an atmospheric light absorber. Only recently have global modeling studies started incorporating
radiative forcing by organic aerosol (OA) absorption (Wang et al., 2014; Saleh et al., 2015; Lin et
al., 2014; Wang et al., 2018). Thus, having accurate estimates for OA absorption is necessary to
help improve climate models.

A convenient and prevalent methodology of measuring OA absorption is based on collecting 53 aerosol particles on a filter substrate followed by extracting the organic compounds into a solvent. 54 This analytical method is used in many studies as it ideally excludes any interference from EC and 55 56 primarily provides the absorption spectra of extracted OC (Mo et al., 2017; Chen and Bond, 2010; 57 Liu et al., 2013). The absorbance of organic chromophores in the solvent extract is measured using 58 an ultraviolet-visible (UV-Vis) spectrophotometer and measured absorbance values can be converted to corresponding solvent phase absorption coefficients $(b_{abs,sol})$. However, this 59 methodology has limitations as it is unable to represent size-dependent absorption properties of 60 the extracted OA (Liu et al., 2013; Washenfelder at al., 2015; Moosmüller et al., 2011). To correct 61 62 for this limitation, the complex refractive index (RI) of OC is estimated by assuming the real part and calculating the imaginary part for extracted OC using babs.sol and dissolved OC concentration, 63 the complex RI is then used along with a number size distribution as inputs to Mie theory for 64 calculating the particle-phase absorption coefficient for dissolved OC. In addition to discrepancies 65 66 between particle and solvent phase optical properties, the method suffers from biases due to 67 incomplete extraction of organics by different solvents (Chen and Bond, 2010; Liu et al., 2013) which lead to differences in values of $b_{abs,sol}$ obtained from different solvents. The significance 68 and extent of this bias varies based on the OC extraction efficiency of a given solvent and would 69

be negligible for solvents extracting 100% of organic chromophores. A combination of inefficient 70 71 organic carbon extraction and the methods inability to measure size-dependent OA absorption 72 properties can result in significant errors to optical properties obtained using this method. Despite the low OC extraction efficiency of water (Chen and Bond, 2010) and large potential for errors, 73 past studies have used light absorption by water soluble organic carbon (WSOC) as a surrogate for 74 75 OA optical properties (Bosch et al., 2014; Kirillova et al., 2014a; Kirillova et al., 2014b). However, the use of water as an OA surrogate is decreasing with more recent studies using methanol to 76 77 extract OC (Cheng et al., 2016; Shen at al., 2017; Xie et al., 2017). While methanol has a higher 78 OC extraction efficiency than water (Chen and Bond, 2010), its efficiency is limited ranging from 85-98% (Cheng et al., 2016; Xie et al., 2017) which can lead to misrepresentation of OA optical 79 properties if the unextracted fraction correspond to extremely low volatility organic carbon 80 (ELVOCs) or similar organic chromophores which have large light absorption efficiencies (Saleh 81 et al., 2014), underscoring the need for a more complete extraction protocol. In addition to 82 83 problems with incomplete OC extraction, previous studies have attempted to correct for sizedependent biases using absorption coefficients determined with Mie theory and provided a narrow 84 range of solvent-dependent scaling factors from 2 for water extracts to 1.8 for methanol extracts, 85 86 all corresponding to a mean particle diameter of 0.5 µm (Liu et al., 2013; Liu et al., 2016; Washenfelder et al., 2015). Sun et al. (2007) performed theoretical calculations and postulated a 87 88 correction range of 0.69 - 0.75 for OC particles with diameters much smaller than the wavelength 89 of light. These correction factors while applicable to these individual systems, might not be 90 extensible to aerosol emissions from other combustion events. However, many studies have used 91 scaling factors from such studies on absorption coefficients obtained from solvent phase optical 92 measurements despite potential differences in system dependent biases for each experiment (Kim

et al., 2016; Zhang et al., 2017; Wang et al., 2018). To the authors knowledge, no attempts have
been made to explicitly study or quantify these biases with varying aerosol intrinsic properties,
such as the EC/OC ratios, and single scattering albedo (SSA), even though these properties have
shown to be well correlated with OA optical properties (Zhang et al., 2013; Saleh at al., 2014;
Bergstrom et al., 2007).

98 In-situ measurement of particulate-phase absorption coefficient is commonly and accurately accomplished using a photoacoustic spectrometer (PAS) (Lack et al., 2006; Arnott et al., 2005; 99 Arnott et al., 2003). However, on its own, a single-wavelength PAS cannot distinguish between 100 absorption by OC and BC aerosol and it typically measures the total particle-phase absorption 101 102 coefficient ($b_{abs,tot}$) of the aerosol population in the cell (Moosmüller et al., 2009). One can make use of a multi-wavelength PAS using which the OA absorption coefficient $(b_{abs,OA})$ could be 103 separated out from that of BC absorption, based on the difference in BC and OA Absorption 104 Ångström Exponent (AÅE) (Washenfelder et al., 2015; Arola et al., 2011; Kirchstetter and 105 Thatcher; 2012). The AAE for pure BC is well-constrained at 1 in the visible and near-infrared 106 107 wavelengths (Moosmüller et al., 2009). The value of $b_{abs,OA}$ is calculated as the difference between $b_{abs,tot}$ and the BC absorption coefficient. A possible technique to measure the bias between 108 particle and solvent phase organic absorption $(b_{abs,OA}/b_{abs,sol})$ can thus be established by carrying 109 out simultaneous measurements of solution- and particle-phase absorption properties during a 110 study. Determining $b_{abs,OA}$ using this method gives large errors when BC absorption coefficient 111 is large or comparable to $b_{abs,tot}$ as $b_{abs,OA}$ would be a small number obtained by the subtraction 112 of two large numbers limiting the use of this technique for relatively low EC/OC ratios. 113

Here, we burnt a range of different biomass fuels under different combustion conditions and the 114 resulting aerosol emissions were passed through various in-situ instruments while simultaneously 115 being collected on quartz-fiber filters. The particle phase absorption coefficient was obtained using 116 integrated photoacoustic-nephelometer spectrometers (IPNs) at wavelengths 375, 405 and 1047 117 nm. Organics collected on quartz-fiber filters were extracted in water, acetone, and methanol, and 118 corresponding $b_{abs,sol}$ values were calculated. These values were compared with corresponding 119 $b_{abs,OA}$, and the change in $b_{abs,OA}/b_{abs,sol}$ with varying single scattering albedo (SSA) values and 120 OC/TC ratios was examined. SSA was parametrized with the OC/TC ratios with trends similar to 121 those observed by Pokhrel et al. (2016). AÅE from spectroscopic data for solution and particle 122 phase measurements were compared, and the Mie Theory based correction factor was also 123 124 investigated for a few samples.

125 **2 Methods:**

126 **2.1 Sample generation and collection**

Fig. 1 is a schematic diagram of our experimental setup, which consisted of a sealed 21 m^3 127 128 stainless-steel combustion chamber housing a fan for mixing and recirculation (Sumlin et al., 129 2018b). Aerosol samples were generated by burning several types of biomass including pine, fir, 130 grass, sage, and cattle dung (details are provided in the Supplementary Information). During a chamber burn, 10-50 g of a given biomass was placed in a stainless-steel pan and ignited by a 131 butane lighter. The chamber exhaust was kept closed for the duration of a given experiment. The 132 133 biomass bed was either allowed to burn to completion or it was prematurely extinguished and 134 brought to a smoldering phase by extinguishing the flame beneath a lid. Different combustion 135 conditions were used to generate samples with varying properties: OC/TC ratios ranged from 0.55-

136 1, and SSA values ranged from 0.56-0.98 for wavelengths of 375, 405, and 1047 nm.

137 For one set of experiments, the particles were directly sampled from the chamber; in another set, 138 the sampling was done from a hood placed over the burning biomass. A diffusion dryer removed 139 excess water from the sample stream, and the gas-phase organics were removed by a pair of 140 activated parallel-plate semi-volatile organic carbon (SVOC) denuders. The gas-phase organics were stripped to reduce artifacts produced by the adsorption of organic vapors on the quartz filters. 141 The aerosols were finally sent to a 208-liter stainless-steel barrel, from which they were 142 continuously sampled by the three IPNs. Some phase repartitioning of condensed SVOC into the 143 144 vapor phase may take place post the denuders in our holding tank and would introduce a positive bias to our filter-based measurements. The experiments were conducted in two sets, the first set 145 included a scanning mobility particle sizer (SMPS, TSI, Inc.) and size measurements from this 146 147 instrument were used in Mie Theory calculations detailed in Section 2.3. The SMPS was not used 148 in the second set of experiments due to problems with aerosol flows in the system. However, the SMPS data from the first set of experiments gave us an estimate of the range over which the size 149 150 distributions varied and was used to obtain the geometric mean of the size distribution. The real-151 time absorption and scattering coefficients were measured by the IPNs, and samples were simultaneously collected on quartz fiber filters once a steady state signal was achieved. The 152 153 absorption and scattering coefficients were used to calculate the SSA, which is simply the scattering coefficient divided by the extinction coefficient. Radiative forcing calculations for 154 155 absorbing OC require good estimates of OC absorption at different SSA values (Lin et al, 2014; Feng et al, 2013; Chakrabarty et al, 2010) underscoring the need to study OA absorption biases as 156 a function of SSA. The particles were passed through the filter samplers at a flowrate of 5 L min⁻ 157

¹, with sampling times ranging from 2-15 minutes. Two or more filters were collected for a given steady state condition. One of these filters was used to determine the OC and EC fractions of the deposited particles, and the other filters were used for the extraction experiments. The only exception to this case is one sample of emissions from dung combustion where the resultant aerosol was assumed to be purely organic based on a purely smoldering combustion phase and by comparing with optical properties from previous experiments.

164 **2.2** Analytical techniques

165 2.2.1 Absorption by solvent extracted OC

Quartz filters (Pallflex Tissuquartz, 47 mm diameter) collected during sampling were split into 166 167 four quarters, and each quarter was extracted using either deionized water, acetone, hexane, or methanol. The absorption by hexane extracts were low and prone to errors, so data for its extracts 168 were not analyzed. The filters were placed in 3-5 ml of the solvent for 24 hours. The filter was not 169 sonicated to reduce artifacts from mechanical dislodging of BC particles (Phillips and Smith, 2017) 170 and to avoid changes in chemical composition caused by acoustic cavitation (Mutzel et al., 2012). 171 172 The solvent volumes were measured both before and after the extraction and the differences between the two measurements were within 8%. The extracts were then passed through syringe 173 filters with 0.22 µm pores to remove any suspended particles introduced during the extraction 174 175 process.

The light absorbance of the extracts was measured using a UV-Vis spectrophotometer (Varian Inc., Cary 50) at wavelengths from 300 nm to 800 nm. To compare the absorbance ($A(\lambda)$) of chromophores in the solution with the absorption coefficient of the particles in the atmosphere, all 179 absorbance values were converted to solution-phase absorption coefficients at given wavelengths 180 $(b_{abs,sol}(\lambda))$ (Liu et al., 2013):

181
$$b_{abs,sol}(\lambda) = (A(\lambda) - A(700)) \frac{V_l}{V_a * l} \cdot \ln(10),$$
 (1)

where V_l is the volume of solvent the filter was extracted into, V_a is the volume of air passed over the given filter area, and l is the optical path length that the beam traveled through the cuvette (1 cm). Absorbance at a given wavelength is normalized to absorbance at 700 nm to account for any signal drift within the instrument. Absorbance at 700 nm was negligible and close to zero for the analyzed samples indicating no absorption at long wavelengths and little to no signal drift for the instrument. The resulting absorption coefficient (m⁻¹) was multiplied by ln(10) to convert from log base 10 (provided by the UV-Vis spectrophotometer) to natural log.

189 2.2.2 Absorption by BC and OC in particle phase

To estimate the BC absorption at 375 nm and 405 nm, the absorption data from the IPN operated in the infrared regime at a wavelength of 1047 nm was converted to equivalent BC particulate absorption at the near UV wavelengths, using a BC absorption Ångström exponent ($AÅE_{BC}$) value of 1 (Kirchstetter et al., 2004; Andreae and Gelencsér, 2006). It was assumed that all the absorption at 1047 nm could be attributed to BC aerosol (Bahadur et al., 2012). The BC light absorption coefficient at shorter wavelengths ($b_{abs,BC}(\lambda)$) was calculated by:

196
$$b_{abs,BC}(\lambda_1) = b_{abs,tot}(1047) \cdot \left(\frac{\lambda_1}{1047}\right)^{-A\mathring{A}E_{BC}},$$
 (2)

197 where λ_1 is the wavelength at which the absorption will be calculated and AAE is defined for a 198 pair of wavelengths λ_1 , λ_2 as the exponent in a power law expressing the ratio of the absorption 199 coefficients as follows (Moosmüller et al., 2009):

200
$$A\mathring{A}E(\lambda_1\lambda_2) = \frac{\ln\left[\frac{b_{abs}(\lambda_1)}{b_{abs}(\lambda_2)}\right]}{\ln\left[\frac{\lambda_2}{\lambda_1}\right]}$$
(3)

AÅ*E* is an optical descriptor of the inherent material property. For BC particles, typical values of AÅ*E* \approx 1, while for OC particles AÅE > 4 (Moosmüller et al., 2009). The value of $b_{abs,BC}$ at 375nm and 405nm was then subtracted from $b_{abs,tot}$ at those wavelengths to calculate $b_{abs,OA}$. The ratio $b_{abs,OA}/b_{abs,sol}$ was calculated to represent the scaling bias between the bulk solvent phase absorption coefficient and OA absorption coefficient.

206 The organic and elemental carbon compositions of the filters were measured with a thermal-optical 207 OC/EC analyzer (Sunset Laboratory, Tigard, OR) using the Interagency Monitoring of Protected Visual Environments (IMPROVE)-A Thermal/Optical Reflectance (TOR) analysis method (Chow 208 209 et al., 2007a). The OC/TC ratios were assumed to be constant for a given steady state IPN reading, which allowed us to relate the absorption data to the OC/TC data. The assumption was tested by 210 performing EC/OC analysis of two filters collected during a given steady state for a burn. The 211 212 OC/TC ratio remained unchanged or within experimental error for the burns and results for the 213 EC/OC analysis of tested filters are provided in Table S1 and S2 of the Supplementary Information.

214 2.2.3 Uncertainty using Monte Carlo simulations

The uncertainties due to error propagation were evaluated using a Monte Carlo approach. The true measurement value was assumed to possess a Gaussian probability distribution with mean and standard deviation values corresponding to measured values and instrument specifications respectively. Calculations were performed by randomly selecting values based on the probability distribution for the different variables and corresponding values for $b_{abs,OA}/b_{abs,sol}$ were estimated. A total of N = 10000 iterations were performed for each data point and each simulation was rerun 100 times till the $b_{abs,OA}/b_{abs,sol}$ value converged for the calculations. The propagated error due to uncertainty in important variables was then calculated as the standard deviation of $b_{abs,OA}/b_{abs,sol}$ values acquired over simulations. A pseudocode for the Monte Carlo calculation is detailed in the Supplementary Information along with Table S4 which denotes typical mean and standard deviation values used for variables with uncertainties.

226 **2.3** Mie theory calculations

A commonly used method to correct for differences between the chromophore absorption in solution and aerosol particle absorption is by using Mie Theory (Liu et al., 2013; Washenfelder et al., 2015). The imaginary part (k) of the complex refractive index m = n + ik can be determined from bulk solution phase absorption data and converted to equivalent OA absorption using Mie Theory along with assumptions regarding the shape of the particles and the real part of the complex refractive index of the particle.

To find *k*, the mass absorption efficiency (α/ρ) was determined using the absorbance data and the OC mass concentration in the solution (Liu et al., 2013):

235
$$\frac{\alpha(\lambda)}{\rho} = \frac{b_{abs,sol}(\lambda)}{M},$$
(4)

where $b_{abs,sol}(\lambda)$ is the solvent-phase absorption coefficient determined in Eq. (1), and M is the mass concentration of OC in the solution. In the given study, the OC mass concentration was measured for some of the water extracts using a total organic carbon (TOC) analyzer (Shimadzu, TOC-L). The water-soluble organic carbon (WSOC) was then used to estimate α/ρ of the solution. The calculated α/ρ was further used to determine *k* for the WSOC by (Chen and Bond 2010):

241
$$k(\lambda) = \frac{\rho \lambda \left(\frac{\alpha(\lambda)}{\rho}\right)}{4\pi},$$
 (5)

where λ is the light wavelength at which k needs to be calculated, and ρ is the density of the 242 243 dissolved organic compounds. A ρ value of 1.6 (Alexander et al., 2008) was used to calculate the k values, and was also used in all subsequent calculations using density. It is important to note that 244 k values obtained using this method will represent optical characteristics of OC mass and not total 245 organic mass. A Mie based inversion algorithm was used to extract the real part of the refractive 246 247 index (n) using data from the SMPS and IPN (Sumlin et al., 2018a). If size distributions extended 248 over the SMPS measurement range, the data were extrapolated using a lognormal equation. A sensitivity analysis was performed by varying the *n* value from 1.4 to 2, and the change in Mie 249 calculated absorption was within 18%. The size distribution for the WSOC was estimated 250 251 assuming the same geometric mean and standard deviation as that of the original aerosol, but with number concentrations calculated based on the extracted mass. Calculations for the number 252 concentration are provided in the Supplementary Information. After the size distribution and 253 complex refractive index were determined, they were used to calculate the absorption coefficient 254 based on Mie Theory, which was then compared to $b_{abs.sol}$ to verify the traditional Mie based 255 scaling factors for converting from solution to particle phase absorption. 256

257 **3 Results and discussion**

258 **3.1** Absorption bias correlated with single scattering albedo

Fig. 2 shows the trends in $b_{abs,OA}/b_{abs,sol}$ for fresh organic aerosol emissions with varying SSA. The different fuel types are marked with distinct markers and the error bars are estimated from the results of the Monte Carlo simulation and account for uncertainties in IPN measurements, UV-Vis

spectrophotometer measurements, filter sampling flowrates, BC AÅE based uncertainties and 262 263 extract volume measurements. Measured SSA for pure fractal BC aggregates have values between 0.1-0.3 (Schnaiter et al., 2003; Bond et al., 2013) depending on the size of the BC monomers 264 265 (Sorensen 2001), and due to this particularly low SSA of BC compared to OC, an increase in BC 266 content of aerosol composition would lead to decreasing SSA. This relationship is explored further 267 in Section 3.2 and 3.3. Fig 2. indicates that the light absorbed by methanol and acetone extracts 268 were almost identical and would imply that the amount and type of OC extracted by the two 269 solvents were similar, as seen in other studies as well (Chen and Bond, 2010; Wang et al., 2014). 270 For some dung samples, the bias for methanol and acetone extracts was close to 0.6 at SSA values 271 of 0.95. These bias values were near to the theoretical prediction of 0.69 - 0.75 by Sun et al. 272 (2007) for particle sizes much smaller than the wavelength of light, even though our size 273 distributions were not significantly smaller than the wavelength of 405 nm. This could indicate that predictions by Sun et al. (2007) are valid for sizes comparable to the wavelength of light as 274 well, but more such observations are necessary to obtain conclusive results. The reason for 275 observed differences in the bias between water and methanol extracts are discussed further in 276 Section 3.3. The differences between the mean values of $b_{abs,OA}/b_{abs,sol}$ at 375 and 405 nm were 277 less than or close to the errors associated with them, hence any trends with wavelength were not 278 279 explored. In addition to this, there were no obvious trends that could be explained using fuel type, leading us to not explore trends with fuel type either. 280

The value of $b_{abs,OA}/b_{abs,sol}$ approached a constant in the measured range of data. A power law ($y = k_0 + k_1 x^{k_2}$) was used to fit the points in Fig. 2, and the corresponding fit parameters, along with root mean square error (RMSE) values, are listed in Table 1. The fit was performed using the curve fitting tool in MATLAB and the RMSE values were calculated in Microsoft Excel. The

power law fits were deficient in capturing the true behavior of the bias with SSA but performed 285 better than corresponding mean values and step function curves. The parametrizations presented 286 in this section are representative of laboratory-based biomass burning (BB) aerosol emissions in 287 this study and are provided to mathematically visualize trends in the data. These parametrizations 288 might not be extensible to other emissions and should not be used for determining OA absorption 289 290 bias in other systems. The contribution of BC absorption coefficient to total absorption increases with larger EC fraction of the aerosol which results in significant errors while extrapolating BC 291 absorption from longer wavelengths. Based on other studies, BC AÅE values range from 0.85 to 292 1.1 (Lack et al., 2008; Bergstrom et al., 2007; Lan et al., 2013). In Fig. 2, for data points below the 293 perforated lines at SSA values smaller than 0.7 at 375 nm and smaller than 0.825 at 405 nm, the 294 errors due to uncertainties is BC AÅE were greater than 30% and are a result of increasing BC 295 296 mass fractions at these SSA values. The large uncertainties at lower SSA values indicate that the method described here is best suited to determine $b_{abs,OA}/b_{abs,sol}$ for particles with relatively 297 higher SSA values. 298

299 3.2 SSA parametrized with OC/TC

300 A linear relationship between the SSA and the EC/TC ratio was observed by Pokhrel et al. (2016). To replicate the linear trends observed by Pokhrel et al., we studied the correlation between SSA 301 and OC/TC ratio (which is simply the EC/TC ratio subtracted from 1). Fig. 3 shows the variation 302 in SSA with change in the OC/TC ratio of the aerosol. The OC/TC ratio was determined using the 303 IMPROVE-A TOR protocol with a thermal optical EC/OC analyzer at Sunset laboratories. The 304 305 data was parametrized using an orthogonal distance regression (ODR) to account for errors in the OC/TC ratio and resulting fits along with data points are plotted in Fig. 3. ODR is different from 306 a standard linear regression as it accounts for errors in both the independent and dependent 307

variables by minimizing least square errors perpendicular to the regression lines rather than vertical 308 errors as in standard linear regression. The ODR fits are linear with RMSE values of 0.04 and 0.02 309 310 for wavelengths 375 nm and 405 nm respectively. In Fig. 3, the points corresponding to high OC/TC ratios are associated with SSA values that are close to 1, because pure OC aerosols are 311 predominantly light scattering. The fit yielded SSA values of 0.89 and 0.96 at 375 and 405 nm 312 313 respectively for pure OA indicating that the fits represent a spectral dependence of absorption which is characteristic of brown carbon optical properties because the SSA values for pure OC are 314 315 below 1 at both wavelengths and SSA at 375 nm is lower than that at 405 nm. (Chakrabarty et al. 2010). 316

A linear relation between the SSA and the EC/TC ratio (which is simply the OC/TC ratio 317 subtracted from 1) was also observed by Pokhrel et al. (2016). However, when the data from that 318 study were converted to OC/TC values for comparison, it was noted that the slopes and intercepts 319 320 of the resulting fits were different from those observed in this study. Table 2 has a list of the slope 321 and intercept of fits for comparable wavelengths in both studies, along with the RMSE for our fit. A likely reason for dissimilar slopes and intercepts between the two studies could be due to 322 323 discrepancies in EC/OC ratios obtained using the same temperature protocol. Inter-comparison 324 studies have shown that different labs using the same sample with identical thermal protocols may produce different results (Panteliadis et al., 2015). The instrument bias could be such that obtained 325 326 OC/TC ratios would have a proportional offset between different instruments leading to similar 327 linear trends but with different slopes which might be the case here. Another plausible reason for 328 the discrepancy could be positive artifacts in EC/OC analysis due to gas phase SVOCs being adsorbed on the quartz surface because of phase partitioning of these compounds in the holding 329 tank. This reason seems less likely due relatively small sampling times for the aerosols. To assess 330

the performance of our parametrizations, we compared our fit to data obtained by Liu et al. (2014) 331 at 405 nm for BB aerosol. Data from the plots were extracted using Web Plot Digitizer (Rohatgi 332 333 2010) and was plotted with our fit in Fig. 4. We observed that our fits predicted SSA well at OC/TC ratios > 0.7 with a RMSE value of 0.06 compared to 0.08 by Pokhrel et al. (2016) but predictions 334 were worse for 405 nm at lower OC/TC ratios as is also evident from the relatively high SSA value 335 336 of 0.39 for pure EC obtained using our parametrization. Most observations for soot SSA are lower than those predicted by our 405 nm parametrizations (Bond et al., 2013, Schnaiter at al., 2003) 337 338 with our projections being closer to SSA observed by Radney et al. (2014). Generally, OC/TC 339 ratios are greater than 0.7 for laboratory and field BB (Xie et al., 2019; Akagi et al., 2011; Zhou et al., 2017; Xie et al., 2017) which reduces concerns about underperformance of our fits for 405 nm 340 at low OC/TC ratios. It would be appropriate to use these parametrizations to determine a 341 reasonable range for SSA values rather than use them as a surrogate to determine actual SSA for 342 a given BB aerosol plume. A modification of Fig. 4 which compares the linear fits by Liu et al. 343 344 (2014) and Pokhrel et al. (2016) with our parametrizations is provided in the Supplementary Information. 345

346 Despite the difference between our fits and those by Pokhrel et al. (2016), a useful conclusion from 347 Fig. 3 is that the OC/TC ratio determined using the IMPROVE-A protocol and SSA of BB aerosol have a linear dependence. This dependence, however, has high variations at OC/TC ratios very 348 349 close to 1, where fuel type and burn conditions dictate the composition and absorption properties (Chen and Bond, 2010; Budisulistiorini et al., 2017) of organics released and hence a larger range 350 351 of SSA values exist at those OC/TC ratios. Further studies need to be conducted using more fuels with a variety of distinct size distributions and burn conditions to determine the validity and exact 352 parameters for the fit. 353

354 **3.3** Absorption bias correlated with OC/TC ratio

Fig. 5 depicts the variation in $b_{abs,OA}/b_{abs,sol}$ for primary OA with different OC/TC ratios. 355 Because the OC/TC ratio and the SSA are well correlated, we expect to see a similar trend for Fig. 356 357 5 as in Fig. 2. Similar to Fig. 2, the bias in Fig. 5 increases with decreasing OC/TC ratio and 358 approaches a constant for the three solvents. A power law like the one in Fig. 2 was fit to the data 359 in Fig. 5. The fit parameters for the different solvents at the two wavelengths, along with the RMSE value for each fit, are presented in Table 3. We reiterate that the parametrizations for 360 $b_{abs,OA}/b_{abs,sol}$ as a function of OC/TC ratio depicted here are applicable to our system and should 361 not be used to calculate the bias in other systems. The exclusivity of depicted fit parameters to our 362 363 system excuses their relatively poor RMSE while representing the bias with OC/TC ratio. The parametrizations are provided to represent some quantitative measure to the data rather than just 364 analyze the trends qualitatively. The large error bars from the Monte Carlo simulations at high EC 365 fractions are mainly due to uncertainties associated with the BC AÅE. At lower OC/TC ratios, the 366 367 contribution of BC absorption to total particle-phase absorption coefficient is more pronounced, 368 leading to high uncertainties while extrapolating the coefficient to shorter wavelengths. It is apparent from Fig. 5 that these errors in the bias are more prominent at OC/TC ratios below 0.75. 369 The burns with relatively high EC fractions are not representative of typical laboratory or field 370 BB. As mentioned earlier, typical laboratory BB have OC/TC ratios > 0.7 (Xie et al., 2017; Akagi 371 372 et al., 2011; Pokhrel et al., 2016; Xie et al., 2019) and > 0.9 for field BB (Aurell et al., 2015; Zhou et al., 2017; Xie et al., 2017). Thus, data presented in Fig. 5 with relatively large errors and EC/TC 373 374 ratios > 0.25 are not representative of typical BB in either laboratory or field settings which may 375 warrant their exclusion from most analysis. We have still included these data points in our plots 376 and Tables but have excluded their use in data analysis due to the high errors associated with them.

In Fig. 5, the difference in magnitude of the bias between methanol/acetone extracts and water 377 extracts increase as EC fraction of the aerosol increases. An increase in the emissions of ELVOCs 378 with increasing EC/OC ratios was observed by Saleh et al. (2014) and we hypothesize that these 379 ELVOCs which have high mass absorption efficiencies (Saleh et al., 2014; Di Lorenzo and Young 380 2016) could have a lower solubility in water than methanol or acetone which would explain the 381 increasing difference in $b_{abs,OA}/b_{abs,sol}$ values between water and methanol/acetone extracts. 382 Some of the generated ELVOCs might be insoluble in methanol and acetone as well which would 383 384 lead to the observed increase in the OA absorption bias with decreasing OC fraction of the aerosol. 385 Based on the observed trends, these ELOCs would not be released indefinitely but tend towards a 386 constant above a given EC/OC fraction, mimicking an exponential behavior comparable to 387 observed trends in wavelength dependence for biomass burning OA with EC/OC ratios (Saleh et 388 al., 2014). This would lead to the bias approaching a constant value (due only to particle size effects) with decreasing OC/TC ratios and in turn the aerosol SSA. Future studies can look at the 389 type and amount of ELVOCs released as a function of the EC/OC ratio of the aerosol and ascertain 390 if their solubility in these solvents is a function of their EC content. 391

392 **3.4 Variations in AÅE with solvents and OC/TC ratios**

The AÅE values, for organics extracted in different solvents and those obtained from $b_{abs,OA}$ are compared in Table 4. The AÅE values along with the errors for OA measurements were calculated between $\lambda = 375$ and 405 nm using the Monte Carlo simulation. The AÅE for OC extracts were calculated using Eq. 3 based on $b_{abs,sol}$ and corresponding errors were propagated based on uncertainties in UV-Vis measurements. Consistent with previous studies (Chen and Bond, 2010; Zhang et al., 2013; Liu et al., 2013), the AÅE values of water extracts were larger than the AÅE of acetone and methanol extracts. Experiments by Zhang et al., (2013) observed that polycyclic 400 aromatic hydrocarbons (PAHs) absorbed light at longer wavelengths close to the visible region. 401 Organic solvents such as methanol have a higher extraction efficiency for these compounds than 402 water leading to higher absorption by methanol extracts at longer wavelengths which results in 403 lower AÅE (Zhang et al., 2013).

The AÅE calculated for OA ranged from 6.9 \pm 1.7 to 15.6 \pm 0.6 (excluding data with OC/TC > 404 0.75) which are slightly larger than AÅE values reported by most studies (Pokhrel et al., 2016; 405 Lewis et al., 2008). However, these studies report AÅE values in the visible range, which might 406 be lower than aerosol AÅE values in the UV range as observed by Chen and Bond (2010) for OA 407 extracts. The range of AÅE observed for water, acetone and methanol extracts were similar to those 408 observed by Chen and Bond (2010). A t-test for data presented in Table 4 shows that AÅE values 409 for OA were greater than their solution phase counterparts for both methanol (N = 17, p = 0.0007) 410 and acetone (N = 17, p = 0.0002). The difference in AÅE of OA and water extracts were statistically 411 insignificant (N = 17, p = 0.25), but these differences were statistically significant at OC/TC ratios 412 ≥ 0.9 (N = 12, p < 0.05) where uncertainties due to BC absorption are lower. The reason for these 413 differences could be a combination of artifacts due to inefficient extraction of organics absorbing 414 light at lower wavelengths and the absence of size dependent absorption in the solvent phase which 415 might not capture effects of enhanced particle phase absorption at lower wavelengths. These bulk 416 solvent measurements of AÅE suggest that they might not be representative of spectral dependence 417 of OC in the particle phase, and future studies and models should be cautious while using AÅE 418 data from solvent-phase measurements to be representative of the particle phase. 419

420 **3.5** Scaling factors based on Mie calculations

To check the reproducibility of the conventionally used correction factor of 2, the absorption coefficient determined from the bulk solvent absorbance using Eq. (1) was compared to absorption coefficients calculated using Mie theory for three samples of smoldering sage. The EC/OC analysis (IMPROVE-A protocol) determined that these samples consisted purely of OC, and because the SMPS measurements and TOC analysis were only performed on the first set of samples, the three samples of sage were considered optimum for Mie calculations.

The Mie based scaling factors for converting solution phase absorption coefficients to particulate 427 428 absorption for the three samples are presented in Table 5. TOC and EC/OC analysis indicated that 429 a similar fraction of organics at $61 \pm 2\%$ were extracted from all three samples. The Mie calculated 430 scaling factors at 375 nm and 405 nm are close to 2 as observed in previous studies (Liu et al., 2013; Washenfelder et al., 2015) indicating that the conventional technique provides reproducible 431 results. The values for these scaling factor vary from 2 to 2.1 at 375 nm and 2.2 to 2.3 at 405 nm. 432 433 However, it is important to note that these scaling factors were not representative of actual biases 434 for determining OA absorption from solution phase as observed in Table 5. Thus, while a Mie based correction factor of 2 can be duplicated, it is not representative of actual biases as also 435 corroborated by observations from Fig. 2 and Fig. 5. We recommend future studies to use caution 436 437 and judgement when using a priori scaling factors for determining OA absorption using solvent extraction techniques. 438

439 **4** Conclusions

Under controlled laboratory conditions, we determined artifacts associated with optical properties
of the solvent phase as compared to particle phase counterparts for fresh OA emissions from
biomass combustion. We combusted a range of different wildland fuels under different combustion
conditions, generating a span of different SSA and OC/TC values. The SSA values ranged from

0.55 to 0.87 at 375 nm, and from 0.69 to 0.95 at 405 nm, the OC/TC values ranged from 0.55 to 444 1. We observed an increasing difference in $b_{abs,OA}/b_{abs,sol}$ for water and methanol extracts with 445 increasing EC fraction of the aerosol. The decrease in absorption by water extracts with decreasing 446 OC/TC ratios was hypothesized to occur due to a decrease in extraction of ELVOC or similar 447 448 compounds with high mass absorption efficiencies by water. We also demonstrated that the SSA 449 and OC/TC ratios can be well parametrized with a linear fit that captures the effects of brown carbon aerosol. We also determined that bulk solvent measurements of AÅE are not representative 450 451 of spectral dependence of OC in the particle phase. Finally, we analyzed the validity and reproducibility of the conventionally used scaling factor of 2 for determining OA absorption 452 453 coefficients from water extracts of organics and noted that, while the factor is reproducible, its use can misrepresent OA absorption coefficients. We recommend that future studies use caution while 454 455 applying a priori scaling factors to their systems as these factors might not be extensible to OA emissions from all combustion processes. A comprehensive technique which improves extraction 456 efficiency with accurate knowledge of particle size distributions is necessary to determine correct 457 458 scaling relations.

For future experiments, a better technique to quantify BC absorption at lower wavelengths, such as a thermodenuder to strip off all OC, or a single particle soot photometer along with core-shell Mie calculations can be used to determine BC absorption and decrease uncertainties related to BC absorption observed during experiments using this technique. Zhang et al. (2013) observed lower *AÅE* for WSOC from a particle into liquid sampler (PILS) than for methanol extracts. The hypothesis was that the highly dilute environment in PILS increased dissolution of organics in water. This suggests that extraction of organics can be increased by heavily diluting the samples. 466 This can be combined with highly accurate spectrometers similar to the technique used by467 Hecobian et al. (2010) to reduce some of the biases due to incomplete OA extraction.

468 Author Contributions

- 469 RKC conceived of this study and designed the experiments. SB and WMH collected the fuels for
- 470 the experiments and performed EC/OC analysis on the sampled filters. NJS and AP carried out the
- 471 experiments. NJS analysed the data and prepared the manuscript with input from all co-authors.

472 Acknowledgements

- 473 This work was partially supported by the National Science Foundation under Grant No.
- 474 AGS1455215, NASA ROSES under Grant No. NNX15AI66G.

475 **References**

- Akagi, S., Yokelson, R. J., Wiedinmyer, C., Alvarado, M., Reid, J., Karl, T., Crounse, J., and
 Wennberg, P.: Emission factors for open and domestic biomass burning for use in atmospheric
 models, Atmos. Chem. Phys., 11, 4039-4072, https://doi.org/10.5194/acp-11-4039-2011, 2011.
- Alexander, D. T., Crozier, P. A., and Anderson, J. R.: Brown carbon spheres in East Asian outflow
 and their optical properties, Science, 321, 833-836, https://doi.org/10.1126/science.1155296,
 2008.
- Andreae, M., and Gelencsér, A.: Black carbon or brown carbon? The nature of light-absorbing
 carbonaceous aerosols, Atmos. Chem. Phys., 6, 3131-3148, https://doi.org/10.5194/acp-6-31312006, 2006.
- Arnott, W., Moosmüller, H., Sheridan, P., Ogren, J., Raspet, R., Slaton, W., Hand, J., Kreidenweis,
 S., and Collett Jr, J.: Photoacoustic and filter-based ambient aerosol light absorption
 measurements: Instrument comparisons and the role of relative humidity, J. Geophys. Res-Atmos.,
 108, AAC 15-1-AAC 15-11, https://doi.org/10.1029/2002JD002165, 2003.
- Arnott, W. P., Hamasha, K., Moosmüller, H., Sheridan, P. J., and Ogren, J. A.: Towards aerosol
 light-absorption measurements with a 7-wavelength aethalometer: Evaluation with a
 photoacoustic instrument and 3-wavelength nephelometer, Aerosol Sci. Tech., 39, 17-29,
 https://doi.org/10.1080/027868290901972, 2005.
- Arola, A., Schuster, G., Myhre, G., Kazadzis, S., Dey, S., and Tripathi, S.: Inferring absorbing
 organic carbon content from AERONET data, Atmos. Chem. Phys., 11, 215-225,
 https://doi.org/10.5194/acp-11-215-2011, 2011.

- Aurell, J., Gullett, B. K., and Tabor, D.: Emissions from southeastern US Grasslands and pine
- 497 savannas: Comparison of aerial and ground field measurements with laboratory burns, Atmos.
 498 Environ., 111, 170-178, https://doi.org/10.1016/j.atmosenv.2015.03.001, 2015.
- Bahadur, R., Praveen, P. S., Xu, Y., and Ramanathan, V.: Solar absorption by elemental and brown
 carbon determined from spectral observations, P. Nat. Acad. Sci., 109 (43), 17366-17371,
 https://doi.org/10.1073/pnas.1205910109, 2012.
- Bergstrom, R. W., Pilewskie, P., Russell, P. B., Redemann, J., Bond, T. C., Quinn, P. K., and
 Sierau, B.: Spectral absorption properties of atmospheric aerosols, Atmos. Chem. Phys., 7, 5937-
- 504 5943, https://doi.org/10.5194/acp-7-5937-2007, 2007.
- 505 Bond, T. C., Doherty, S. J., Fahey, D., Forster, P., Berntsen, T., DeAngelo, B., Flanner, M., Ghan,
- 506 S., Kärcher, B., and Koch, D.: Bounding the role of black carbon in the climate system: A scientific
- 507 assessment, J. Geophys. Res-Atmos., 118, 5380-5552 , https://doi:10.1002/jgrd.50171, 2013.
- 508 Bosch, C., Andersson, A., Kirillova, E. N., Budhavant, K., Tiwari, S., Praveen, P., Russell, L. M.,
- Beres, N. D., Ramanathan, V., and Gustafsson, Ö.: Source-diagnostic dual-isotope composition
- 510 and optical properties of water-soluble organic carbon and elemental carbon in the South Asian
- outflow intercepted over the Indian Ocean, J. Geophys. Res-Atmos., 119, 11,743-711,759,
- 512 https://doi:10.1002/2014JD022127, 2014.
- 513 Budisulistiorini, S. H., Riva, M., Williams, M., Chen, J., Itoh, M., Surratt, J. D., and Kuwata, M.: 514 Light-absorbing brown carbon aerosol constituents from combustion of Indonesian peat and
- 515 biomass, Environ. Sci. Technol., 51, 4415-4423, https://doi.org/10.1021/acs.est.7b00397, 2017.
- Chakrabarty, R., Moosmüller, H., Chen, L.-W., Lewis, K., Arnott, W., Mazzoleni, C., Dubey, M.,
 Wold, C., Hao, W., and Kreidenweis, S.: Brown carbon in tar balls from smoldering biomass
 combustion, Atmos. Chem. Phys., 10, 6363-6370, https://doi.org/10.5194/acp-10-6363-2010,
 2010.
- 520 Chen, Y., and Bond, T.: Light absorption by organic carbon from wood combustion, Atmos. Chem.
 521 Phys., 10, 1773-1787, https://doi.org/10.5194/acp-10-1773-2010, 2010.
- 522 Cheng, Y., He, K.-b., Du, Z.-y., Engling, G., Liu, J.-m., Ma, Y.-l., Zheng, M., and Weber, R. J.:
 523 The characteristics of brown carbon aerosol during winter in Beijing, Atmos. Environ., 127, 355-
- 524 364, https://doi.org/10.1016/j.atmosenv.2015.12.035, 2016.
- Chow, J. C., Watson, J. G., Chen, L.-W. A., Chang, M. O., Robinson, N. F., Trimble, D., and Kohl,
 S.: The IMPROVE_A temperature protocol for thermal/optical carbon analysis: maintaining
 consistency with a long-term database, J. Air Waste Manage., 57, 1014-1023,
 https://doi.org/10.3155/1047-3289.57.9.1014, 2007a.
- 529 Chow, J. C., Yu, J. Z., Watson, J. G., Hang Ho, S. S., Bohannan, T. L., Hays, M. D., and Fung, K. K.: The application of thermal methods for determining chemical composition of carbonaceous 530 531 aerosols: А review. J. Environ. Sci. Heal. А, 42, 1521-1541, https://doi.org/10.1080/10934520701513365, 2007b. 532
- Di Lorenzo, R. A., and Young, C. J.: Size separation method for absorption characterization in
 brown carbon: Application to an aged biomass burning sample, Geophys. Res. Lett., 43, 458-465,
 https://doi.org/10.1002/2015GL066954. 2016.

- 536 Feng, Y., Ramanathan, V., and Kotamarthi, V.: Brown carbon: a significant atmospheric absorber
- of solar radiation?, Atmos. Chem. Phys., 13, 8607-8621, https://doi.org/10.5194/acp-13-86072013, 2013.
- Hecobian, A., Zhang, X., Zheng, M., Frank, N., Edgerton, E. S., and Weber, R. J.: Water-Soluble 539 540 Organic Aerosol material and the light-absorption characteristics of aqueous extracts measured the over Southeastern United States, Atmos. Chem. Phys., 10. 5965-5977, 541 https://doi.org/10.5194/acp-10-5965-2010, 2010. 542
- Husain, L., Dutkiewicz, V. A., Khan, A., and Ghauri, B. M.: Characterization of carbonaceous
 aerosols in urban air, Atmos. Environ., 41, 6872-6883,
 https://doi.org/10.1016/j.atmosenv.2007.04.037, 2007.
- IPCC: Climate Change: The Physical Science Basis, Contribution of Working Group I to the UN
 IPCC's 5th Assessment Report, Cambridge University Press, New York, USA, 2013.
- Kim, H., Kim, J. Y., Jin, H. C., Lee, J. Y., and Lee, S. P.: Seasonal variations in the light-absorbing
 properties of water-soluble and insoluble organic aerosols in Seoul, Korea, Atmos. Environ., 129,
 234-242, https://doi.org/10.1016/j.atmosenv.2016.01.042, 2016.
- Kirchstetter, T. W., Novakov, T., and Hobbs, P. V.: Evidence that the spectral dependence of light
 absorption by aerosols is affected by organic carbon, J. Geophys. Res-Atmos., 109, D21208
 https://doi.org/10.1029/2004JD004999, 2004.
- Kirchstetter, T., and Thatcher, T.: Contribution of organic carbon to wood smoke particulate matter
 absorption of solar radiation, Atmos. Chem. Phys., 12, 6067-6072, https://doi.org/10.5194/acp-126067-2012, 2012.
- Kirillova, E. N., Andersson, A., Han, J., Lee, M., and Gustafsson, Ö.: Sources and light absorption
 of water-soluble organic carbon aerosols in the outflow from northern China, Atmos. Chem. Phys.,
 14, 1413-1422, https://doi.org/10.5194/acp-14-1413-2014, 2014a.
- Kirillova, E. N., Andersson, A., Tiwari, S., Srivastava, A. K., Bisht, D. S., and Gustafsson, Ö.:
 Water-soluble organic carbon aerosols during a full New Delhi winter: Isotope-based source
 apportionment and optical properties, J. Geophys. Res-Atmos., 119, 3476-3485,
 https://doi:10.1002/2013JD020041, 2014b.
- Lack, D. A., Lovejoy, E. R., Baynard, T., Pettersson, A., and Ravishankara, A.: Aerosol absorption
 measurement using photoacoustic spectroscopy: Sensitivity, calibration, and uncertainty
 developments, Aerosol Sci. Tech., 40, 697-708, https://doi.org/10.1080/02786820600803917,
 2006.
- Lack, D. A., Cappa, C. D., Covert, D. S., Baynard, T., Massoli, P., Sierau, B., Bates, T. S., Quinn,
 P. K., Lovejoy, E. R., and Ravishankara, A.: Bias in filter-based aerosol light absorption
 measurements due to organic aerosol loading: Evidence from ambient measurements, Aerosol Sci.
 Tech., 42, 1033-1041, https://doi.org/10.1080/02786820802389277, 2008.

- 572 Lan, Z.-J., Huang, X.-F., Yu, K.-Y., Sun, T.-L., Zeng, L.-W., and Hu, M.: Light absorption of
- black carbon aerosol and its enhancement by mixing state in an urban atmosphere in South China,
 Atmos. Environ., 69, 118-123, https://doi.org/10.1016/j.atmosenv.2012.12.009, 2013.
- Lin, G., Penner, J. E., Flanner, M. G., Sillman, S., Xu, L., and Zhou, C.: Radiative forcing of
 organic aerosol in the atmosphere and on snow: Effects of SOA and brown carbon, J. Geophys.
 Res-Atmos., 119, 7453-7476, https://doi.org/10.1002/2013JD021186, 2014.
- 578 Liu, J., Bergin, M., Guo, H., King, L., Kotra, N., Edgerton, E., and Weber, R.: Size-resolved
- measurements of brown carbon in water and methanol extracts and estimates of their contribution
 to ambient fine-particle light absorption, Atmos. Chem. Phys., 13, 12389-12404,
 https://doi.org/10.5194/acp-13-12389-2013, 2013.
- Liu, S., Aiken, A. C., Arata, C., Dubey, M. K., Stockwell, C. E., Yokelson, R. J., Stone, E. A.,
- Jayarathne, T., Robinson, A. L., and DeMott, P. J.: Aerosol single scattering albedo dependence
- on biomass combustion efficiency: Laboratory and field studies, Geophys. Res. Lett., 41, 742-748,
 https://doi.org/10.1002/2013GL058392, 2014.
- Liu, J., Lin, P., Laskin, A., Laskin, J., Kathmann, S. M., Wise, M., Caylor, R., Imholt, F.,
 Selimovic, V., and Shilling, J. E.: Optical properties and aging of light-absorbing secondary
 organic aerosol, Atmos. Chem. Phys., 16, 12815-12827, https://doi.org/10.5194/acp-16-128152016, 2016.
- 590 Mo, Y., Li, J., Liu, J., Zhong, G., Cheng, Z., Tian, C., Chen, Y., and Zhang, G.: The influence of
- solvent and pH on determination of the light absorption properties of water-soluble brown carbon,
 Atmos. Environ., 161, 90-98, https://doi.org/10.1016/j.atmosenv.2017.04.037, 2017.
- Moosmüller, H., Chakrabarty, R., and Arnott, W.: Aerosol light absorption and its measurement:
 A review, J. Quant. Spectrosc. Ra., 110, 844-878, https://doi.org/10.1016/j.jqsrt.2009.02.035,
 2009.
- Moosmüller, H., Chakrabarty, R., Ehlers, K., and Arnott, W.: Absorption Ångström coefficient,
 brown carbon, and aerosols: basic concepts, bulk matter, and spherical particles, Atmos. Chem.
 Phys., 11, 1217-1225, https://doi.org/10.5194/acp-11-1217-2011, 2011.
- Mutzel, A., Rodigast, M., Iinuma, Y., Böge, O., and Herrmann, H.: An improved method for the
 quantification of SOA bound peroxides, Atmos. Environ., 67, 365-369,
 https://doi.org/10.1016/j.atmosenv.2012.11.012, 2013.
- Panteliadis, P., Hafkenscheid, T., Cary, B., Diapouli, E., Fischer, A., Favez, O., Quincey, P.,
 Viana, M., Hitzenberger, R., Vecchi, R., Saraga, D., Sciare, J., Jaffrezo, J. L., John, A., Schwarz,
 J., Giannoni, M., Novak, J., Karanasiou, A., Fermo, P., and Maenhaut, W.: ECOC comparison
 exercise with identical thermal protocols after temperature offset correction instrument
 diagnostics by in-depth evaluation of operational parameters, Atmos. Meas. Tech., 8, 779-792,
 https://doi.org/10.5194/amt-8-779-2015, 2015.

- Phillips, S. M., and Smith, G. D.: Spectroscopic comparison of water-and methanol-soluble brown
 carbon particulate matter, Aerosol Sci. Tech., 51, 1113-1121,
 https://doi.org/10.1080/02786826.2017.1334109, 2017.
- 611 Pokhrel, R. P., Wagner, N. L., Langridge, J. M., Lack, D. A., Jayarathne, T., Stone, E. A.,
- 612 Stockwell, C. E., Yokelson, R. J., and Murphy, S. M.: Parameterization of single-scattering albedo
- 613 (SSA) and absorption Ångström exponent (AAE) with EC/OC for aerosol emissions from biomass
- 614 burning, Atmos. Chem. Phys., 16, 9549-9561, https://doi.org/10.5194/acp-16-9549-2016, 2016.
- Radney, J. G., You, R., Ma, X., Conny, J. M., Zachariah, M. R., Hodges, J. T., and Zangmeister,
- 616 C. D.: Dependence of soot optical properties on particle morphology: measurements and model
- 617 comparisons, Environ. Sci. Technol., 48, 3169-3176, https://doi.org/10.1021/es4041804, 2014.
- Ramanathan, V., and Carmichael, G.: Global and regional climate changes due to black carbon,
 Nat. Geosci., 1, 221–227, https://doi.org/10.1038/ngeo156, 2008.
- Rohatgi, A. :WebPlotDigitalizer: HTML5 based online tool to extract numerical data from plot
 images. URL http://arohatgi.info/WebPlotDigitizer/app, 2012.
- 622 Saleh, R., Robinson, E. S., Tkacik, D. S., Ahern, A. T., Liu, S., Aiken, A. C., Sullivan, R. C., Presto, A. A., Dubey, M. K., and Yokelson, R. J.: Brownness of organics in aerosols from biomass 623 624 burning linked to their black carbon content, Nat. Geosci., 7, 647-650, https://doi.org/10.1038/ngeo2220, 2014. 625
- 626 Saleh, R., Marks, M., Heo, J., Adams, P. J., Donahue, N. M., and Robinson, A. L.: Contribution of brown carbon and lensing to the direct radiative effect of carbonaceous aerosols from biomass 627 628 biofuel burning emissions, J. Geophys. Res-Atmos., 120, 10,285-10,296, and 629 https://doi:10.1002/2015JD023697., 2015.
- Schnaiter, M., Horvath, H., Möhler, O., Naumann, K.-H., Saathoff, H., and Schöck, O.: UV-VISNIR spectral optical properties of soot and soot-containing aerosols, J. Aerosol Sci., 34, 14211444, https://doi.org/10.1016/S0021-8502(03)00361-6, 2003.
- Shen, Z., Lei, Y., Zhang, L., Zhang, Q., Zeng, Y., Tao, J., Zhu, C., Cao, J., Xu, H., and Liu, S.:
 Methanol extracted brown carbon in PM 2.5 over Xi'an, China: seasonal variation of optical
 properties and sources identification, Aerosol Sci. Eng., 1, 57-65, https://doi.org/10.1007/s41810017-0007-z, 2017.
- 637 Sorensen, C.: Light scattering by fractal aggregates: a review, Aerosol Sci. Tech., 35, 648-687,
 638 https://doi.org/10.1080/02786820117868, 2001.
- 639 Sumlin, B. J., Heinson, W. R., and Chakrabarty, R. K.: Retrieving the aerosol complex refractive
- 640 index using PyMieScatt: A Mie computational package with visualization capabilities, J. Quant.
- 641 Spectrosc. Ra., 205, 127-134, https://doi.org/10.1016/j.jqsrt.2017.10.012, 2018a.

- Sumlin, B. J., Heinson, Y. W., Shetty, N., Pandey, A., Pattison, R. S., Baker, S., Hao, W. M., and
 Chakrabarty, R. K.: UV–Vis–IR spectral complex refractive indices and optical properties of
 brown carbon aerosol from biomass burning, J. Quant. Spectrosc. Ra., 206, 392-398,
 https://doi.org/10.1016/j.jqsrt.2017.12.009, 2018b.
- Sun, H., Biedermann, L., and Bond, T. C.: Color of brown carbon: A model for ultraviolet and
 visible light absorption by organic carbon aerosol, Geophys. Res. Lett., 34, L17813,
 https://doi.org/10.1029/2007GL029797, 2007.
- 649 Wang, X., Heald, C., Ridley, D., Schwarz, J., Spackman, J., Perring, A., Coe, H., Liu, D., and
- Clarke, A.: Exploiting simultaneous observational constraints on mass and absorption to estimate
 the global direct radiative forcing of black carbon and brown carbon, Atmos. Chem. Phys., 14,
 10989-11010, https://doi.org/10.5194/acp-14-10989-2014, 2014.
- 052 10707 11010, https://doi.org/10.5174/dep-14-10707 2014, 2014.
- Wang, X., Heald, C. L., Liu, J., Weber, R. J., Campuzano-Jost, P., Jimenez, J. L., Schwarz, J. P.,
- and Perring, A. E.: Exploring the observational constraints on the simulation of brown carbon,
- 655 Atmos. Chem. Phys., 18, 635-653, 10.5194/acp-18-635-2018, https://doi.org/10.5194/acp-18-635-
- 656 2018, 2018.
- Washenfelder, R., Attwood, A., Brock, C., Guo, H., Xu, L., Weber, R., Ng, N., Allen, H., Ayres,
 B., and Baumann, K.: Biomass burning dominates brown carbon absorption in the rural
 southeastern United States, Geophys. Res. Lett., 42, 653-664,
 https://doi.org/10.1002/2014GL062444, 2015.
- Kie, M., Hays, M. D., and Holder, A. L.: Light-absorbing organic carbon from prescribed and
 laboratory biomass burning and gasoline vehicle emissions, Sci. Rep., 7, 7318,
 https://doi.org/10.1038/s41598-017-06981-8, 2017.
- Kie, M., Chen, X., Hays, M. D., and Holder, A. L.: Composition and light absorption of Ncontaining aromatic compounds in organic aerosols from laboratory biomass burning, Atmos.
 Chem. Phys., 19, 2899-2915, https://doi.org/10.5194/acp-19-2899-2019, 2019.
- Yang, M., Howell, S., Zhuang, J., and Huebert, B.: Attribution of aerosol light absorption to black
 carbon, brown carbon, and dust in China–interpretations of atmospheric measurements during
 EAST-AIRE, Atmos. Chem. Phys., 9, 2035-2050, https://doi.org/10.5194/acp-9-2035-2009, 2009.
- Zhang, X., Wang, Y., Zhang, X., Guo, W., and Gong, S.: Carbonaceous aerosol composition over
 various regions of China during 2006, J. Geophys. Res-Atmos., 113, D14111,
 https://doi.org/10.1029/2007JD009525, 2008.
- Zhang, X., Lin, Y.-H., Surratt, J. D., and Weber, R. J.: Sources, composition and absorption
 Ångstrom exponent of light-absorbing organic components in aerosol extracts from the Los
 Angeles Basin, Environ. Sci. Technol., 47, 3685-3693, https://doi.org/10.1021/es305047b, 2013.

- Zhang, Y., Forrister, H., Liu, J., Dibb, J., Anderson, B., Schwarz, J. P., Perring, A. E., Jimenez, J.
 L., Campuzano-Jost, P., and Wang, Y.: Top-of-atmosphere radiative forcing affected by brown
 carbon in the upper troposphere, Nat. Geosci., 10, 486–489, https://doi.org/10.1038/ngeo2960,
 2017.
- Kang, Zhou, Y., Xing, X., Lang, J., Chen, D., Cheng, S., Lin, W., Xiao, W., and Liu, C.: A comprehensive
- biomass burning emission inventory with high spatial and temporal resolution in China, Atmos.
- 682 Chem. Phys., 17, 2839-2864, https://doi.org/10.5194/acp-17-2839-2017, 2017.

Figures and Tables:



Fig. 1: A schematic representing the experimental setup. The aerosol emissions were either
sampled directly from the chamber wall or through a hood placed directly above the combusting
biomass.



Fig. 2: Variation in $b_{abs,OA}/b_{abs,sol}$ with change in the SSA at (a) 375 nm and (b) 405 nm (N = 21). The error bars represent one standard deviation from the mean and were calculated using Monte Carlo simulations. The black markers represent water extracts, red markers represent acetone extracts and blue markers represent methanol extracts. The perforated lines separate points at lower SSA, which have high errors greater than 30% due to uncertainties in BC AÅE, from the data at high SSA.



Fig. 3: SSA at (a) 375 nm and (b) 405 nm as a function of the OC/TC ratio (N= 49). The solid red lines are ODR fits to the data and the solid blue lines represent the 95% confidence intervals. The errors in OC/TC ratios were determined by the quadrature sum of uncertainties from EC/OC analysis and the error in SSA were negligible



Fig. 4: Measured SSA values by Liu et al. (2014) for controlled laboratory combustion
experiments plotted with the solid red line representing the ODR parametrization determined in
this study.



Fig. 5: The values of $b_{abs,OA}/b_{abs,sol}$ plotted with the OC/TC ratio, instead of the SSA, as in Fig.

2. Black markers represent data for water extracts, red markers represent data for acetone extracts,

and blue markers represent data for methanol extracts.

| | Wavelength | Solvent Fit Parameters | | | | |
|-------------------------------|------------|------------------------|----------------|----------------|----------------|------|
| | (nm) | | k ₀ | k ₁ | k ₂ | RMSE |
| | 375 | Water | 10.1 (±2.1) | -39.8 (±177.1) | 16.1 (±31.3) | 2.2 |
| 1 | | Acetone | 4.2 (±0.8) | -117.4 (±36.9) | 27.5 (±37.5) | 1.1 |
| $b_{abs,OA}$ $b_{abs,sol}$ | | Methanol | 4.2 (±0.8) | -69.1 (±451.6) | 25.8 (±45.4) | 1.1 |
| | 405 | Water | 14 (±4.2) | -42.4 (±70.5) | 27.4 (±35.5) | 2.6 |
| | | Acetone | 3.9 (±1.1) | -95.6 (±609.9) | 68.3 (±121.8) | 1.3 |
| | | Methanol | 4.8 (±1.4) | -49.1 (±250.3) | 53.1 (±98) | 1.5 |
| | | | | | | |

716 Table 1: Fit coefficients for $b_{abs,OA}/b_{abs,sol}$ as a function of SSA $(y = k0 + k1 (SSA)^{k_2})$ for

tested solvents and the fuels analyzed in this study along with the RMSE value for each fit.

719 Table 2. ODR regression coefficients along with errors in brackets for plots of SSA v/s OC/TC

ratios (y = m (OC/TC) + c) for the different biomass fuels used in this study, and parameters for

| 721 | ODR fit from Pokhrel | et al. (2016) for 4 | 405 nm, along with | RMSE values for | r our fits. |
|-----|----------------------|---------------------|--------------------|-----------------|-------------|
| | | | | | |

| | Wavelength (nm) | m | С | RMSE |
|------------|-----------------|---------------|---------------|------|
| This study | 375 | 0.71 (±0.04) | 0.18 (±0.03) | 0.04 |
| | 405 | 0.57 (± 0.02) | 0.39 (± 0.02) | 0.02 |
| Pokhrel | 405 | 1.07 (±0.04) | -0.13 (±0.04) | ~~ |
| | | | | |

Table 3: Fit parameters for ratios of the absorption coefficient of organics in the particle phase to the absorption coefficient of the solvent phase, as a function of the OC/TC ratio ($y = k_0 + k_1(OC/TC)^{k_2}$) for the fuels analyzed in this study, along with the RMSE value for each fit.

| | Wavelength | Solvent Fit Parameters | | | | |
|----------------------------------|------------|------------------------|----------------|----------------|----------------|------|
| | (nm) | | k ₀ | k ₁ | k ₂ | RMSE |
| | 375 | Water | 10.7 (±1.8) | -5.1 (±2.5) | 25 (±36) | 2.0 |
| | | Acetone | 4.1 (±0.8) | -2.4 (±1.2) | 58 (±86.2) | 1.0 |
| $rac{b_{abs,OA}}{b_{abs,bulk}}$ | | Methanol | 4.1 (±0.7) | -1.8 (±1.1) | 71.5 (±139.6) | 0.9 |
| | 405 | Water | 13.2 (±2.4) | -9.5 (±3.2) | 20.8 (±20.9) | 2.5 |
| | | Acetone | 4.1 (±0.9) | -3.1 (±1.3) | 43.3 (±49.6) | 1.0 |
| | | Methanol | 4.8 (±1.1) | -3.2 (±1.7) | 49 (±69.5) | 1.3 |

| Fuel | | | AÅE _{375–405} | | | | |
|-------|------|----------------|------------------------|---------------|---------------|--|--|
| ruei | | OA | Water | Acetone | Methanol | | |
| Dung | 1 | 13.7 ± 2.3 | 8.0 ± 2.0 | 5.3 ± 1.4 | 5.2 ± 1.3 | | |
| | 1 | 15.3 ± 2.4 | 9.0 ± 2.0 | 5.9 ± 0.4 | 7.8 ± 0.6 | | |
| | 1 | 15.6 ± 0.6 | 7.5 ± 1.8 | 4.6 ± 0.3 | 4.5 ± 0.9 | | |
| | 1 | 14.9 ± 2.7 | 8.6 ± 1.2 | 5.3 ± 0.2 | 6.8 ± 0.4 | | |
| Sage | 1 | 13.9 ± 1.9 | 10.9 ± 1.2 | 8.6 ± 0.7 | 8.8 ± 1.1 | | |
| | 1 | 10.7 ± 1.5 | 10.7 ± 4.5 | 6.3 ± 3.2 | 7.3 ± 2.9 | | |
| | 0.97 | 10.6 ± 2.4 | 9.9 ± 1.4 | 5.2 ± 0.8 | 5.8 ± 0.7 | | |
| | 0.79 | 7.4 ± 2.9 | 12.3 ± 2.4 | 8.6 ± 0.8 | 9.2 ± 1.2 | | |
| | 0.79 | 8.2 ± 2.4 | 10.6 ± 2.2 | 8.7 ± 0.8 | 8.3 ± 1.3 | | |
| | 0.71 | 10.4 ± 1.4 | 7.5 ± 3.1 | 6.3 ± 1.7 | 6.4 ± 2.1 | | |
| | 0.55 | 9.9 ± 4.2 | 6.5 ± 4.8 | 3.8 ± 2.0 | 3.6 ± 2.8 | | |
| Grass | 0.99 | 10.1 ± 2.4 | 12.1 ± 4.6 | 7.8 ± 0.9 | 7.5 ± 1.2 | | |
| | 0.78 | 9.9 ± 3.1 | 10.2 ± 2.3 | 8.5 ± 0.5 | 9.6 ± 0.6 | | |
| | 0.78 | 6.9 ± 1.7 | 9.7 ± 3.8 | 7.5 ± 0.8 | 7.3 ± 1.6 | | |
| | 0.77 | 9.0 ± 4.0 | 8.2 ± 1.6 | 8.1 ± 0.6 | 8.4 ± 0.9 | | |
| Pine | 0.98 | 11.8 ± 1.0 | 9.4 ± 2.0 | 8.6 ± 0.8 | 8.1 ± 1.1 | | |
| | 0.98 | 8.7 ± 1.9 | 9.6 ± 3.4 | 8.4 ± 1.8 | 8.6 ± 1.5 | | |
| | 0.95 | 14.2 ± 3.5 | 16.4 ± 1.3 | 11.8 ± 0.9 | 12.8 ± 1.3 | | |
| | 0.9 | 8.2 ± 2.4 | 9.1 ± 2.3 | 8.8 ± 1.6 | 8.7 ± 2.0 | | |
| | 0.7 | 16 ± 10.9 | 9.9 ± 3.1 | 6.3 ± 2.2 | 5.8 ± 2.1 | | |
| _ | 0.6 | 17.4 ± 10.8 | 6.4 ± 3.3 | 5.2 ± 2.8 | 5.4 ± 2.9 | | |

Table 4: The AÅE of OA from various fuels extracted in water, acetone, and methanol, along with

728 the $A \text{\AA} E$ calculated for $b_{abs,OA}$.

| | | Geometric st mean (in nm) de | Geometric | Mie based Sc | aling Factor | IPN based bias | |
|-----|------|---------------------------------|--------------------|--------------|--------------|----------------|---------------|
| | Fuel | | standard deviation | 375 nm | 405 nm | 375 nm | 405 nm |
| | Sage | 397 | 1.3 | 2.0 ± 0.4 | 2.3 ± 0.4 | 2.6 ± 0.6 | 1.6 ± 0.6 |
| | | 271 | 1.32 | 2.1 ± 0.4 | 2.3 ± 0.4 | 2.8 ± 0.6 | 1.9 ± 0.3 |
| | | 159 | 1.59 | 2.0 ± 0.4 | 2.2 ± 0.4 | 2.8 ± 0.5 | 1.8 ± 0.4 |
| 734 | | | | | | | |
| 735 | | | | | | | |
| 736 | | | | | | | |

Table 5: Correction factors for bulk solution absorption to particle phase absorption, based on MieTheory calculations.