



# Evaluation and uncertainty investigation of the NO<sub>2</sub>, CO and NH<sub>3</sub> modeling over China under the framework of MICS-Asia III

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Abstract. Despite significant progress in improving chemical transport models (CTMs), applications of these modeling endeavours are still subject to the large and complex model uncertainty. Model Inter-Comparison Study for Asia III (MICS-

- 35 Asia III) provides the opportunity to assess the capability and uncertainty of current CTMs in East Asia applications. In this study, we evaluated the multi-model simulations of nitrogen dioxide (NO<sub>2</sub>), carbon monoxide (CO) and ammonia (NH<sub>3</sub>) over China under the framework of MICS-Asia III. Compared with MICS-Asia II, modeling results were provided by a larger number of independent groups from different countries/regions and covered a longer period of time (one-full year). Furthermore, most of these groups used a common emission inventory, common meteorological inputs, and the same modeling
- 40 domain and horizontal resolutions. New observations over North China Plain (NCP) and Peral River Delta (PRD) were also





available in MICS-Asia III, allowing model evaluations over highly industrialized regions. The results show that most models well captured the monthly and spatial patterns of  $NO_2$  in NCP though  $NO_2$  levels were slightly underestimated. Relatively poor performance in  $NO_2$  simulations was found in PRD with larger root mean square error and lower spatial correlation coefficients, possibly due to the relative coarse model resolutions. All models significantly underpredicted CO concentrations both in NCP

- 5 and PRD, with annual mean concentrations 65.4% and 61.4% underestimated by the ensemble mean respectively. Such large underestimations suggest that CO emissions might be underestimated in current emission inventory. In contrast to the good skills in simulating the monthly variations of NO<sub>2</sub> and CO concentrations, all models failed to reproduce the observed monthly variations of NH<sub>3</sub> concentrations in NCP. Most models mismatched the observed peak in July and showed negative correlation coefficients with observations, which may be closely related to the uncertainty in the monthly variations of NH<sub>3</sub> emissions and
- 10 the NH<sub>3</sub> gas-aerosol partitioning. Finally, inter-comparisons of these model results were conducted to quantify the impacts of model uncertainty on the simulations of these gases, which are shown increase with the reactivity of species. Models contained more uncertainty in the NH<sub>3</sub> simulations. This suggests that for some highly active and/or short-lived primary pollutants, like NH<sub>3</sub>, model uncertainty can also take a great part in the forecast uncertainty in addition to the emission uncertainty. Based on these results, some recommendations are made for future studies.

#### 15 1 Introduction

As the rapid growth in East Asia's economy with surging energy consumption and emissions, air pollution has become an increasingly important scientific topic and political concern in East Asia due to its significant environmental and health effects (Lelieveld et al., 2015). Chemical transport models (CTMs), serving as a critical tool both in scientific research and policy makings, have been applied into various air quality issues, including air quality prediction, long-range transport of atmospheric pollutants, development of emission control strategies and understanding of observed chemical phenomena. However, challenges still remain in the air quality modeling due to the multi-scale and non-linear nature of the complex atmospheric processes (Carmichael et al., 2008). Air quality modeling still suffers from large uncertainties related to missing or poorly parameterized physical and chemical processes, inaccurate and/or incomplete emission inventories as well as poorly represented initial and boundary conditions (Carmichael et al., 2008). Understanding such uncertainties and their impacts on

25 air quality modeling is of great importance in assessing the robustness of models for their applications in scientific research and operational use.

There are specific techniques to assess these uncertainties. Monte Carlo simulations, based on different values of model parameters or input fields sampled from a predefined probability density function (PDF), can provide an approximation to the PDF of possible model output and serves as an excellent characterization of the uncertainties of simulations (Hanna et al.,

2001). However, this method is more suited to deal with the uncertainty related to the continuous variables, such as input data or parameters in parameterization. The ensemble method, based on a set of different models, is an alternative approach to accounting for the range of uncertainties (Galmarini et al., 2004;Mallet and Sportisse, 2006). In Europe and North America,





the Air Quality Model Evaluation International Initiative (AQMEII) has been implemented to investigate the model uncertainties of their regional-scale model predictions (Rao et al., 2011). The Model Inter-Comparison Study for Asia (MICS-Asia) provides the opportunity to assess the model performances and uncertainties in East Asia applications. The first Phase of MICS-Asia (MICS-Asia I) was initiated in 1998 and carried out during period 1998–2002, mainly focusing on the long-range transport and depositions of sulfur in Asia (Carmichael et al., 2002). In 2003, the second phase (MICS-Asia II) was

- 5 range transport and depositions of sulfur in Asia (Carmichael et al., 2002). In 2003, the second phase (MICS-Asia II) was initiated and took account of more species related to the regional health and ecosystem protection, including nitrogen compounds, O<sub>3</sub> and aerosols. Launched in 2010, MICS-Asia III greatly expanded the study scope by covering three individual and interrelated topics: (1) evaluate strength and weaknesses of current multi-scale air quality models and provide techniques to reduce uncertainty in Asia; (2) develop a reliable anthropogenic emission inventories in Asia and understanding uncertainty
- 10 of bottom-up emission inventories in Asia; (3) provide multi-model estimates of radiative forcing and sensitivity analysis of short-lived climate pollutants.

This study addresses one component of topic 1, mainly focusing on the three gas pollutants of NO<sub>2</sub>, CO and NH<sub>3</sub>. Compared with MICS-Asia II, more modeling results (fourteen different models with thirteen regional models and one global model) were brought together within the topic 1 of MICS-Asia III, each run by independent modeling groups in China, Japan,

- 15 Korea, United States of America and other countries/regions. The different models contain differences in their numerical approximations (mesh size, time step, chemical solver, etc.) and parameterizations, which represent a sampling of uncertainties residing in the air quality modeling. It would be difficult to interpret the results from intercomparison studies when the models are operated using different meteorological fields and emission inventories. Thus, in MICS-Asia III the models were constrained to be operated under the same conditions by using a common emission inventory, meteorological fields, modeling
- 20 domain and horizontal resolutions. The simulations were also extended from the four months in MICS-Asia II to one-full year of 2010.

NO<sub>2</sub>, CO and NH<sub>3</sub> are three important primary gas pollutants that has wide impacts on the atmospheric chemistry. As a major precursor of O<sub>3</sub>, NO<sub>2</sub> plays an important role in the tropospheric O<sub>3</sub> chemistry, and can also lead to rainwater acidification and secondary aerosol formations by forming nitric acid (HNO<sub>3</sub>) through oxidations (Dentener and Crutzen, 1993;Evans and

- 25 Jacob, 2005). CO is a colorless and toxic gas ubiquitous throughout the atmosphere which is of interest as an indirect greenhouse gas (Gillenwater, 2008) and a precursor for tropospheric O<sub>3</sub> (Steinfeld, 1998). Being the major sink of OH, CO also controls the atmosphere's oxidizing capacity (Levy, 1971; Novelli et al., 1998). As the only primary alkaline gas in the atmosphere, NH<sub>3</sub> is closely associated with the acidity of precipitations for one thing, for another it could react with sulfuric acid and nitric acid forming ammonium sulfate and ammonium nitrate which account for a large proportion of the inorganic
- 30 constituents of the fine particulate matter (Sun et al., 2012;Sun et al., 2013). Assessing their model performances is thus important to help us better understand their environmental consequences and also help explain the model performances for their related secondary air pollutants, such as  $O_3$  and fine particulate matter.

In previous phase of MICS-Asia, no specific evaluation and inter-comparison work has been conducted for these gases, especially for CO and  $NH_3$ . In MICS-Asia II, model performance of  $NO_2$  was evaluated as a relevant species to  $O_3$ (Han, Z. et





al., 2008), however such evaluations were limited to the observation sites from EANET (Acid Deposition Monitoring Network in East Asia). Model evaluations and inter-comparisons in industrialized regions of China has not been performed due to the limited number of monitoring sites in China from EANET, which hindered our understanding of the model performance in industrialized regions. More densely observations over highly industrialized regions of China, namely North China (NCP) and

- 5 Peral River Delta (PRD), were first included in MICS-Asia III, allowing the model evaluations over highly industrialized regions. Meanwhile, emission inventories of these three gases are still subject to large uncertainties (Kurokawa et al., 2013;Li et al., 2017), which is a major source of uncertainties in air quality modeling and forecast. Evaluating these gases' emission inventories from a model perspective is a useful way to identify the uncertainties in emission inventory (Han, K. M. et al., 2008;Noije et al., 2006;Stein et al., 2014;Uno et al., 2007).
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This study is aimed at the evaluation of the NO<sub>2</sub>, CO and NH<sub>3</sub> simulations using the multi-model data from MICS-Asia III. Three questions are addressed: (1) what is the performance of current CTMs in simulating NO<sub>2</sub>, CO and NH<sub>3</sub> over highly industrialized regions of China; (2) what are the potential factors responsible for the model deviations from observations and differences among models; and (3) how large are the model uncertainties on the simulations of these gases.

#### 2 Intercomparison frameworks

# 15 2.1 Description on participating models and input data

Table 1 summarizes the chemical transport models participating in the MICS-Asia III as well as their configurations. These models included NAQPMS (Wang et al., 2001), two version of CMAQ (Byun and Schere, 2006), WRF-Chem (Grell et al., 2005), NU-WRF (Tao et al., 2013; Peters-Lidard et al., 2015), NHM-Chem (Kajino et al., 2018) and GEOS-Chem (http://acmg.seas.harvard.edu/geos/). All models employed the same modeling domain (Fig. 1) and horizontal resolutions (45km×45km) except M13 (0.5° of latitude×0.667° of longitude) and M14 (64km×64km).

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Standard model input datasets of raw meteorological fields, emission inventory and boundary conditions were provided by MICS-Asia III for all participants. Raw meteorological fields were generated from a whole year simulations of 2010 using Weather Research and Forecasting Model (WRF) version 3.4.1 (Skamarock, 2008) with 45km×45km horizontal resolution in a lambert conform projection and vertically 40 layers from surface to the model top (10hPa). Initial and lateral boundary

- 25 conditions (IC & BC) for meteorological simulation were generated every six hours using 1°×1° NCEP FNL (Final) Operational Global Analysis data (ds083.2). Real-time, global, sea surface temperature (RTG\_SST\_HR) analysis were used to generate and update lower BC for sea areas. Four-dimensional data assimilation nudging (Gridded FDDA & SFDDA) was performed during the simulation to increase the accuracy of WRF after the Objective Analysis with NCEP FNL (Final) Operational Global Analysis data (ds083.2), NCEP ADP Global Surface Observation Weather Data (ds461.0) and NCEP ADP
- 30 Global Upper Air and Surface Weather Data (ds337.0).

Standard emission inventory provided by MICS-Asia III were used by all participants. The anthropogenic emissions were provided by a newly developed anthropogenic emission inventory for Asia (MIX) which integrated five national or regional





inventories, including REAS (Regional Emission inventory in Asia) inventory for Asia developed at the Japan National Institute for Environment Studies (NIES), the Multi-resolution Emission Inventory for China (MEIC) developed at Tsinghua University, the high-resolution ammonia emission inventory in China developed at Peking University, the Indian emission inventory developed at Argonne National Laboratory in the United States and the CAPSS (Clean Air Policy Support System)

- 5 Korean emission inventory developed at Konkuk University (Li et al., 2017). Hourly biogenic emissions for the entire year of 2010 in MICS-Asia III were provided by the Model of Emissions of Gases and Aerosols from Nature version 2.04 (MEGANv2.04; Guenther et al., 2006). The Global Fire Emissions Database 3 (GFEDv3; Randerson et al., 2013) was employed for biomass burning emissions. Volcanic SO<sub>2</sub> emissions were provided by the Asia Center for Air Pollution Research (ACAP) with a daily temporal resolution. Air and ship emissions with an annual resolution were provided by the HTAPv2
- 10 emission inventory for 2010 (Janssens-Maenhout et al., 2015). NMVOC emissions were spectated into the model-ready inputs for three chemical mechanisms: CBMZ, CB05 and SAPRC-99 and the weekly and diurnal profiles for emissions were also provided.

MICS-Asia III provides two sets of top and lateral boundary conditions of the modeling domain for the year 2010, which were derived from the 3-hourly global CTM outputs of CHASER (Sudo et al., 2002a; Sudo et al., 2002b) and GEOS-Chem

15 (http://acmg.seas.harvard.edu/geos/) run by Nagoya University (Japan) and the University of Tennessee (USA), respectively. GEOS-Chem was run with 2.5°×2° resolution and 47 vertical layers while CHASER model was run with 2.8°×2.8° and 32 vertical layers.

All participants were required to use the standard model input data to drive their model run so that impacts of model input data on simulations could be minimized. However, models are quite different from each other, and it is difficult to keep all the

- 20 inputs the same. The majority of models have applied the standard meteorology fields, while some other models utilized their own meteorology models, including GEOS-Chem and RAMS-CMAQ. GEOS-Chem was driven by the GEOS-5 assimilated meteorological fields from the Goddard Earth Observing System of the NASA Global Modeling Assimilation Office, and The RAMS-CMAQ was driven by meteorological fields provided by Regional Atmospheric Modeling System (RAMS; Pielke et al., 1992). WRF-Chem utilized the same meteorology model (WRF) as the standard meteorological simulation, but two of
- 25 them considered the two-way coupling effects of pollutants and meteorological fields. The CTM part of NHM-Chem is coupled with the JMA's non-hydrostatic meteorological model (NHM; Saito et al., 2006), but an interface to convert a meteorological model output of WRF to a CTM input was implemented (Kajino et al., 2018). Thus, the standard meteorology field was used in the NHM-Chem simulation, too.

# 2.2 Data and statistical methods

30 All modeling groups performed a base year simulations of 2010 and were required to submit their modeling results according to the data protocol designed in MICS-Asia III. Gridded monthly concentrations of NO<sub>2</sub>, CO, NH<sub>3</sub> and ammonium in the surface layer were used in this study. Note that modeling results from M3 and NH<sub>3</sub> simulations from M8 were excluded due to their incredible results, therefore only thirteen modeling results were actually used in this study.





Hourly observed concentrations of  $NO_2$  and CO were collected over the NCP (19 stations) and PRD (13 stations) from the air quality network over North China (Tang et al., 2012) and the Pearl River Delta regional air quality monitoring network (PRD RAQMN), respectively. The air quality monitoring network over North China was set up by the Chinese Ecosystem Research Network (CERN), the Institute of Atmospheric Physics (IAP) and the Chinese Academy of Sciences (CAS) since 2009 within an area of 500×500 km<sup>2</sup> in northern China. All monitoring stations were selected and set up according to the US

- EPA method designations (Ji et al., 2012). The PRD RAQMN network was jointly established by the government of the Guangdong Province and the Hong Kong Special Administrative Region, consisting of 16 automatic air quality monitoring stations across the PRD (Zhong et al., 2013). Thirteen of these stations are operated by the Environmental Monitoring Centers in the Guangdong Province (this study used) while the other three located in Hong Kong (not included in this study) are
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managed by the Hong Kong Environmental Protection Department. Monthly averaged observations were then calculated for comparison with the simulated monthly surface NO<sub>2</sub> and CO concentrations.

NH<sub>3</sub> observations for long-term period are indeed challenging and limited due to its strong temporal and spatial variability, quick conversion of NH<sub>3</sub> from one phase to another and also its stickiness to the observational instruments (von Bobrutzki et al., 2010). Measurements of surface NH<sub>3</sub> concentrations in year 2010 were not available in this study, however, one-year

- 15 measurement of monthly NH<sub>3</sub> concentrations over China was conducted by the Ammonia Monitoring Network in China (AMoN-China) during September of 2015 to August of 2016 (Pan et al., 2018), which was used as a reference dataset in this study. The AMoN-China was established based on the CERN and the Regional Atmospheric Deposition Observation Network in North China Plain (READ-NCP; Pan et al., 2012), consisting of 53 sites over the whole China with monthly integrated measurements using the passive diffusive technique. 11 stations located in the NCP were used in this study. Distributions of
- 20 the observation stations of NO<sub>2</sub>, CO and NH<sub>3</sub> over NCP and PRD as well as their total emissions in year 2010 provided by MICS-Asia III are presented in Fig. 1.

Mean bias error (MBE), normalized mean bias (NMB), root mean square error (RMSE) and correlation coefficient (R) were calculated for the assessment of model performances. Standard deviation of ensemble models was calculated to measure the ensemble spread and the impacts of model uncertainty. Coefficient of variation (hereinafter, CV), defined as the standard

25 deviation divided by average with larger value denoting lower consistency among models, was also used to measure the impacts of model uncertainty in relative sense. However, by this definition, there is a tendency that lower concentrations are more likely associated with higher value of CV, thus we did not calculate the values of CV over model grids whose simulated concentrations were lower than 0.1 ppbv for NO<sub>2</sub> and NH<sub>3</sub>, and 0.1 ppmv for CO respectively. March–May, Jun–August, September–November and December–February were used to define the season spring, summer, autumn and winter,

30 respectively.





#### **3** Results

#### **3.1** Evaluating the ensemble models with observations

To facilitate comparisons, the modeling results were interpolated to the observation stations by taking the value from the grid cell where the monitoring stations located. Model evaluation metrics defined in Sect. 2.2 were then calculated to evaluate the modeling results against the observations.

#### 3.1.1 NO<sub>2</sub>

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Figure 2 displays the comparisons between the observed and simulated annual mean NO<sub>2</sub> concentrations sampled from different observation sites over NCP (2a) and PRD(2b) with calculated model evaluation metrics summarized in Table 2. M13 is not included in the evaluation of NO<sub>2</sub> since it submitted NO<sub>x</sub> rather than NO<sub>2</sub>. In general, the majority of models 10 underpredicted NO<sub>2</sub> levels both in NCP and PRD. Calculated MBE (NMB) ranges from -6.54 ppbv (-28.4%) to -2.45 (-10.6%) ppbv over NCP and from -9.84 ppbv (-44.0%) to -1.84 ppbv (-8.2%) over PRD among the negatively-biased models. These underpredicted NO<sub>2</sub> concentrations may help explain the overpredicted O<sub>3</sub> concentrations by these models found in the companion MICS paper by Li, J. et al., 2019.  $O_3$  productions can either increase with NO<sub>x</sub> under NO<sub>x</sub> limited conditions or decrease under the NO<sub>x</sub> saturated (also called volatile organic compounds (VOCs) limited) conditions (Sillman, 1999). Both NCP and PRD are industrialized regions in China with high NO<sub>x</sub> emissions (Fig. 1). Observations also showed that NCP and 15 PRD are falling into or changing into the  $NO_x$  saturated regimes (Shao et al., 2009; Jin and Holloway, 2015). Therefore, the underestimated NO<sub>2</sub> concentrations may contribute to the overprediction in O<sub>3</sub> concentrations over these two regions and more details about the O<sub>3</sub> predictions can be found in the companion paper by Li, J. et al., 2019. M5, M8, M9 and M11 in NCP and M5, M8 and M11 in PRD were exceptions that overpredicted NO<sub>2</sub> concentrations. M11 showed good performances in predicting NO<sub>2</sub> levels in NCP with smallest RMSE, while M9 significantly overestimated NO<sub>2</sub> with largest MBE and RMSE. 20

Simulated NO<sub>2</sub> by M8 in PRD was close to the observations with smallest RMSE. Meanwhile, we also found that models exhibited better NO<sub>2</sub> modeling skills in NCP than that in PRD with smaller bias and RMSE.

According to the spatial correlation coefficients (Table 2), all models well reproduced the main features of the spatial variability of  $NO_2$  concentrations in NCP with correlation coefficients ranging from 0.57 to 0.70. However, models failed in capturing the spatial variability in PRD with correlation coefficients only ranged from 0.00 to 0.38. Such low correlation might

be partly attributed to the coarser model resolution (45km) than that of the observations. Some local impacts on the  $NO_2$  concentrations may not be well resolved by the model with 45km horizontal resolution, indicating models' weaknesses in representing the sub-grid processes.

Figure 3 presents the monthly timeseries of the observed and simulated regional mean NO<sub>2</sub> concentrations over NCP (3a) and PRD (3b) from January to December in 2010. The models well captured the monthly variations of NO<sub>2</sub> concentrations both in NCP and PRD. According to Table 2, the correlation coefficient ranges from 0.28 to 0.96 in NCP and from 0.52 to 0.95 in PRD. M8 showed the largest overestimation among all models in summer that MBE (NMB) can reach as 12.1 ppbv





(75.8%), which may help explain the low correlation of this model. M9 exhibited a significant overestimation in winter in NCP with MBE (NMB) up to 22.0 ppbv (79.3%) while much less overestimation or even underestimation (summer) in other seasons. This discrepancy may be explained by that M9 was an online coupled model considering the two-way coupling effects between meteorology and chemistry. During the period with heavy haze, the radiation can be largely reduced by aerosol dimming

5 effects, leading to weakened photochemistry, lowered boundary layer height and the increase of NO<sub>2</sub> concentrations. Severe haze was reported to occur in North China in January 2010, with maximum hourly PM<sub>2.5</sub> concentration even reached as high as ~500 µg/m<sup>3</sup> in urban Beijing (Gao et al., 2018). Such high aerosol loadings in atmosphere could trigger interactions between chemistry and meteorology. Interestingly, M9 did not overestimate NO<sub>2</sub> during winter in PRD as in NCP. This might be related to the lower aerosol concentrations and weaker chemistry-and-meteorology coupling effects in PRD.

#### 10 3.1.2 CO

Similar analyses were performed for CO modeling. All models significantly underestimated the annual mean CO concentrations both in NCP and PRD (Figs. 2c-d and Table 2). Calculated MBE (NMB) ranges from -1.69 ppmv (-76.2%) to -1.16 ppmv (-52.0%) in NCP and from -0.67 ppmv (-69.6%) to -0.50 ppmv (-52.3%) in PRD (Table 2). Such large biases in all the models are probably related to the bias in the CO emission inventory over China. Tang et al., 2013 estimated the CO

- 15 emissions over Beijing and surrounding areas in the summer of 2010 using the ensemble Kalman filter and reported a significant underestimation in the *a priori* estimate of CO emissions. Over the latest decades, global models also reported CO underestimations in north hemisphere (Stein et al., 2014) and a number of global model inversion studies have been conducted to derive the optimized CO emissions, of which most studies reported a significant underestimation of CO emissions in their *a priori* estimates(Bergamaschi et al., 2000;Miyazaki et al., 2012;Petron et al., 2002;Petron et al., 2004). Our findings agree
- 20 with these studies and indicate that more accurate CO emissions are needed in future studies. Model performance in simulating spatial variability of CO concentrations was still poor in PRD according to Table 2 that most models showed negative correlation coefficients.

Timeseries of the observed and simulated regional mean CO concentrations in NCP and PRD (Fig.3c-d) show that the models well reproduced the monthly variation of CO concentrations in NCP and PRD with high temporal correlation

25 coefficients, except M5 in NCP (Table 2). All models, however, underestimated CO concentrations throughout the whole year and showed largest underestimations in winter that MBE (NMB) by ensemble mean can be up -2.1 ppmv (-64.9%) in NCP and -0.75 ppmv (-60.6%) in PRD.

#### 3.1.3 NH<sub>3</sub>

Figure 2e show the observed and simulated annual mean NH<sub>3</sub> concentrations over different sites in NCP. The NH<sub>3</sub> 30 observations in 2010 were not available and the observed monthly mean NH<sub>3</sub> concentrations in NCP from September 2015 to August 2016 were used as a reference data in this study. Negative biases are expected according to the increasing trend of atmospheric ammonia during period 2003–2016 detected by recently retrievals from the Atmospheric Infrared Sounder (AIRS)





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aboard NASA's Aqua satellite (Warner et al., 2016; Warner et al., 2017). Due to the interannual uncertainty, we mainly focused on the disparities among different models rather than the deviation from observations.

Large differences can be seen in simulated NH<sub>3</sub> concentrations from different models. M14 simulated very low concentrations and exhibited the largest negative biases with MBE (NMB) -12.2 ppbv (-66.3%). This may be related to the higher conversion rate of NH<sub>3</sub> to NH<sub>4</sub><sup>+</sup> in M14 which will be discussed in later part of this section. On the contrary, M9 provided much higher NH<sub>3</sub> concentrations with MBE (NMB) up to 21.8 ppbv (118.7%). For the CMAQ models, M1 and M2 exhibited higher NH<sub>3</sub> concentrations and larger spatial variability compared to other CMAQ models. Such discrepancy could

be explained by that M1 and M2 are two model runs using CMAQ v5.0.2. The bi-directional exchange of NH<sub>3</sub> has been

integrated into CMAQ from version 5.0. This module can simulate the emitted and deposited processes of NH<sub>3</sub> between
 atmosphere and the surfaces, allowing the additional NH<sub>3</sub> emissions to the atmosphere (US EPA Office of Research and Development).

As can be seen in Table 2, the observed spatial variations of  $NH_3$  over NCP can be well reproduced by all models (R = 0.57-0.71), indicating that the spatial variations of current  $NH_3$  emissions over NCP are reasonable. However, all models failed to capture the observed monthly variations of  $NH_3$  concentrations with most models mismatching the observed  $NH_3$  peak (July)

- 15 and showing negative correlation coefficients except the M10 and M13 showing good temporal correlations of 0.64 and 0.65, respectively (Fig. 3e and Table 2). This is quite different from the model behavior in simulating the monthly variation of NO<sub>2</sub> and CO concentrations. As can be seen in Fig. 3e, the observation showed the peak concentrations of NH<sub>3</sub> in summer months and low concentrations in autumn and winter, which is consistent with the previous NH<sub>3</sub> observations in NCP (Shen et al., 2011;Xu et al., 2016;Meng et al., 2011). Newly derived satellite-measured NH<sub>3</sub> at 918 hPa averaged between September 2002
- 20 and August 2015 also demonstrated higher concentrations in spring and summer and lower concentrations in autumn and winter (Warner et al., 2016). However, all models predicted a peak concentration in November except M10 in August in and M13 in June.

Figure 4 presents the timeseries of NH<sub>3</sub> emission rates in NCP from January to December in 2010. The monthly variations of NH<sub>3</sub> concentrations simulated by most models were closely related to the variations of the NH<sub>3</sub> emissions. The simulated

25 NH<sub>3</sub> concentrations had three peaks in June, August and November but exhibited a significant decrease in July, which was corresponding to the peaks and the drop of the NH<sub>3</sub> emission rates in these months. The strong relationship between the simulated NH<sub>3</sub> concentrations and emission rates suggests that the poor performance in reproducing the observed monthly variations of NH<sub>3</sub> concentrations is probably related to the uncertainties in the estimated monthly variations of NH<sub>3</sub> emissions by current emission inventories. This is consistent with the recent reconciled bottom-up and top-down estimates of agriculture

30 ammonia emission in China by Zhang et al., 2018.

It is worth noting that the NH<sub>3</sub> simulations suggest that there are important uncertainties in the models beyond those due to emission uncertainty. In order to investigate this issue, the impact of the modeling of the NH<sub>3</sub> gas-aerosol partitioning on the NH<sub>3</sub> simulation was analyzed. Figure 5 displays the timeseries of the simulated total ammonium (NH<sub>x</sub> = NH<sub>3</sub> + NH<sub>4</sub><sup>+</sup>) concentrations in atmosphere along with the ratio of gaseous NH<sub>3</sub> to the total ammonium. Large discrepancy can be seen in





the simulated gas-aerosol partitioning of  $NH_3$  from different models. Among the models, M7 and M9 had higher  $NH_3/NH_x$  ratio than other models, which means that the two models tended to retain more emitted  $NH_3$  in gas phase and had higher  $NH_3$  concentrations than other models. For example, M7 predicted comparable magnitude of total ammonium with most models, while gas  $NH_3$  concentration in M7 accounted for more than 60% of total ammonium in summer and even 90% in winter and

- 5 therefore was higher than most models. On the contrary, M14 showed a much lower NH<sub>3</sub>/NH<sub>x</sub> ratio and produced much lower NH<sub>3</sub> concentrations than most models. Moreover, most models showed lower NH<sub>3</sub>/NH<sub>x</sub> ratio in summer than other seasons, suggesting higher conversion rates of NH<sub>3</sub> from gas phase to aerosol phase in summer. This would be related to the higher yield of ammonium sulfate due to the enhanced photochemical oxidation activity in summer (Husain and Dutkiewicz, 1990) estimated by most models. However, different from the above model results, the NH<sub>3</sub> and NH<sub>4</sub><sup>+</sup>observations over NCP by
- Shen et al., 2011;Xu et al., 2016 showed a lower  $NH_3/NH_x$  ratio and higher ammonium in autumn and wither. Although observed  $NH_4^+$  was largest in summer at a rural site in Beijing, observed  $NH_3/NH_x$  ratio was still highest in summer by Meng et al., 2011. These results indicate that large uncertainties exist in modeling the seasonal variations of gaseous  $NH_3$  and aerosol  $NH_4^+$  partitioning. The formations of  $NH_4^+$  mainly depends on the acid gas concentrations, temperature, water availability (Khoder, 2002) and the flux rates of  $NH_3$  (Nemitz et al., 2001). Compared with spring and summer, the lower temperature and
- 15 higher SO<sub>2</sub> and NO<sub>x</sub> emissions should favor the gas-to-particle phase conversion and lead to higher NH<sub>4</sub><sup>+</sup> aerosol concentrations. This indicate that some reaction pathways of acid productions (H<sub>2</sub>SO<sub>4</sub> or HNO<sub>3</sub>) may be missed in current models, such as aqueous-phase and heterogeneous chemistry (Cheng et al., 2016;Wang et al., 2016;Zheng et al., 2015). Such uncertainty may be another important factor contributing to the misrepresented monthly variation of NH<sub>3</sub> concentrations in NCP by the models.

#### 20 **3.2** Quantifying the impacts of model uncertainty

In this section, we further investigate the differences between the ensemble models to quantify the impacts of model uncertainty on these gases' simulations. As we mentioned in Sect. 2, most of these models employed the same raw meteorology fields and emission inventories over China under the same modeling domain and horizontal resolutions, which composed an appropriate set for investigation of the model uncertainties.

- Figures 6-8 present the simulated annual mean concentrations of NO<sub>2</sub>, CO and NH<sub>3</sub> from thirteen modeling results. Spatial distributions of the simulated NO<sub>2</sub>, CO and NH<sub>3</sub> from different models agreed well with each other, similar to the spatial distributions of their emissions as shown in Fig. 1. High NO<sub>2</sub> concentrations were mainly located in the north and central-east China and several hot-spots of NO<sub>2</sub> could also been detected in the northeast China and PRD by all models. M5, M8, M9, and M11 produced higher NO<sub>2</sub> concentrations than other models especially for M8 which also predicted very high concentrations
- 30 over southeast China. Similar to NO<sub>2</sub>, high CO concentrations were generally located over the north and central-east China as well as the east of Sichuan basin. M8, M9 and M11 produced higher CO concentrations as well. In terms of NH<sub>3</sub>, although most models shared similar spatial patterns of NH<sub>3</sub> simulations, the simulated NH<sub>3</sub> concentrations varied largely from different models. High NH<sub>3</sub> concentrations were mainly located over the north China and India peninsula, which was in accordance





with the distribution of agricultural activity intensity over East Asia. Among these models, M9 and M10 produced much higher NH<sub>3</sub> concentrations over East Asia while M4, M5, M6, M13 and M14 produced much lower concentrations.

Figure 9 gives us more quantitative analysis of the impacts of model uncertainty on the simulations of NH<sub>3</sub> (9a), CO (9b) and  $NO_2$  (9c), denoted by the spatial distribution of the standard deviation (ensemble spread), as well as the corresponding

- 5 distributions of CV for each gas on the annual and seasonal basis. Note that M13 and M14 are excluded in the calculation of ensemble spread and CV to reduce the influences of the meteorological conditions and mesh sizes. It seems that the impacts of model uncertainty increase with the reactivity of the gases.  $NH_3$  simulations were affected most by the model uncertainty, while CO suffered least from the uncertainty in the complex chemistry modeling.
- The ensemble spread of NH<sub>3</sub> simulations exhibited a strong spatial variability with higher value mainly located in NCP. 10 Standard deviation of the annual mean NH<sub>3</sub> concentrations can be over 20 ppby in Henan province and 15 ppby in the south of Hebei province, which is about 60–80% of the ensemble mean in Henan province and 40–60% in Hebei Province according to the CV distribution. As we mentioned in Sect. 3.1.3, these large modeling differences can be partly explained by the different treatments of the bi-directional exchange and gas-aerosol partitioning of NH<sub>3</sub> in different models. We can also see a strong seasonal pattern in the modeling differences of NH<sub>3</sub> in NCP. The ensemble spread was smallest in spring while largest in
- 15 autumn, up to 25 ppbv in most areas of NCP. However, in the relative sense, the modeling differences were larger in summer and winter while less in spring and autumn. The southeast China shared similar magnitude of the ensemble spread which was about 2–5 ppbv and showed weaker seasonal variability. However, the modeling differences were larger than that in NCP in the relative sense that CV can be over 1.0 in all seasons except Summer. This can be due to that the simulated concentrations may be more influenced by the model processes over the areas with low emissions while more constrained by the local emissions over the high emission rate area. 20

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CO was least affected by the model uncertainty among the three gases which is consistent with its weaker chemical activity and longer lifetime in the atmosphere. The ensemble spread of annual mean CO concentration was about 0.05–0.2 ppmv in the east China, only about 20%-30% of the ensemble mean. Meanwhile, CO modeling differences was more uniformly distributed in east China than NH<sub>3</sub> with CV less than 0.3 over most areas of east China. However, large modeling differences were visible over Myanmar during spring when there were high CO emissions from biomass burning. Modeling differences turned to be larger during winter in NCP that the ensemble spread was about 0.3–0.5 ppmv and CV about 0.3–0.4, but were still less than the modeling differences of NH<sub>3</sub>.

 $NO_2$  was mediumly affected by the model uncertainty among the three gases. Ensemble spread of annual mean  $NO_2$ concentration was 5–7.5 ppbv in NCP and 2.5–5 ppbv in the southeast China, which accounted for about 20%–30% of the 30 ensemble mean in former but more than 70% in latter. The ensemble spread was largest in winter which was over 10 ppbv in NCP (30%–40%) and 5–7.5 ppbv in southeast China (over 70%). Similar to NH<sub>3</sub>, southeast China exhibited more modeling differences than NCP in the relative sense that CV can be over 0.7 in most areas of southeast China.

These analysis gives us a quantified result of the impact of model uncertainty on the forecast uncertainty of these gases. It indicated that for some highly active and/or short-lived primary pollutants, like NH<sub>3</sub>, model uncertainty can also take a great





part in the forecast uncertainty besides the emission uncertainty, and further air quality data assimilation and model improvement efforts should pay special attention to such uncertainties.

#### 4 Summary

- In this study, thirteen modeling results of surface NO<sub>2</sub>, CO and NH<sub>3</sub> concentrations from MICS-Asia III has been compared with each other and evaluated against the observations over NCP and PRD. Three questions are trying to be addressed which are related to the performance of current CTMs in simulating NO<sub>2</sub>, CO and NH<sub>3</sub> over highly industrialized regions of China, potential factors responsible for the model deviations from observations and differences among models, as well as the impacts of model uncertainty on these gases' simulations.
- Most models underpredicted NO<sub>2</sub> concentrations in NCP and PRD, which could be an important potential factor contributing to the overpredicted O<sub>3</sub> concentrations in these regions. NO<sub>2</sub> prediction skills were shown to be better in NCP than that in PRD with smaller biases and RMSE. Most models well reproduced the observed temporal and spatial variations of NO<sub>2</sub> concentrations in NCP, while relatively poor performance has been detected in PRD in terms of the spatial variations. All models significantly underestimated the CO concentrations both in NCP and PRD throughout the whole year. Such larger underestimations suggest that CO emissions were very likely to be underestimated in current emission inventories. More
- 15 accurate estimate of CO emissions is needed in future studies. In contrast to the good skills in monthly variations of NO<sub>2</sub> and CO concentrations, a noteworthy phenomenon has been detected that all models failed to reproduce the observed monthly variation of NH<sub>3</sub> concentrations in NCP. Most models mismatched the observed peak and showed negative correlation coefficient, which may be closely related to the uncertainty in monthly variations of NH<sub>3</sub> emissions and the NH<sub>3</sub> gas-aerosol partitioning. Several potential factors were found to be responsible for the model deviation and differences, including
- 20 chemistry-and-meteorology coupling effects, model resolutions, bi-directional exchange of NH<sub>3</sub> and gas-aerosol partition of NH<sub>3</sub>, which could be the important aspects with respect to the model improvements in future. Inter-comparisons of the ensemble model quantified the impacts of model uncertainty on these gases' simulations. It shows that the impacts of model uncertainty increase with the reactivity of these gases. Models contained more uncertainties in the prediction of NH<sub>3</sub> than the other two gases. Based on these findings, some recommendations are made for future studies:
- 25 1) More accurate estimation of CO and NH<sub>3</sub> emissions are needed in future studies. Both bottom-up and top-down method (inversion technique) can help address this problem. For top-down method, advanced inversion techniques can fully take advantage of the CTMs and observations from various platforms (ground stations, satellite and radar, etc.) which has been widely used in the emission inventory estimate and correction (Miyazaki et al., 2017). However, the accuracy of emission inversion is highly subject to the model uncertainties. According to this study, NH<sub>3</sub> simulation was largely affected by the
- 30 model uncertainties, which may cause the overcorrection of emission inventories and reduce the accuracy of emission estimate by inversion work. How to represent the model uncertainties in the current framework of emission inversion is an important aspect in future studies. Things could be better for CO considering its small and weakly spatial-dependent model uncertainties.





2) For some highly active and/or short-lived primary pollutants, like NH<sub>3</sub>, model uncertainty can also take a great part in the forecast uncertainty. Emission uncertainty alone may not be sufficient to explain the forecast uncertainty and may cause underdispersive, and overconfident forecasts. Future studies are also needed in how to better represent the model uncertainties in the model predictions to obtain a better forecast skill.

3) Gas-aerosol partition of NH<sub>3</sub> is shown to be an important source of uncertainties in NH<sub>3</sub> simulation. The formation of NH<sub>4</sub><sup>+</sup> particles is mainly limited by the availability of H<sub>2</sub>SO<sub>4</sub> and HNO<sub>3</sub> under ammonia-rich conditions, which involves very complex chemical reactions, including gas-phase, aqueous-phase and heterogeneous chemistry (Cheng et al., 2016;Wang et al., 2016;Zheng et al., 2015). These processes are needed to be verified and incorporated into models to better represent the chemistry in the atmosphere.

#### 10 Competing interests

The authors declare that they have no conflict of interest.

#### Author contribution

XT, JZ, ZW and GC conducted the design of this study. JF, XW, SI, KY, TN, HL, CK, CL, LC, MZ, ZT, JL, MK, HL, BG contributed to the modelling data. ML and QW provided the emission data. KS provided the CHASER output for boundary

15 conditions. YW, YP, GT provided the observation data. LK and XT performed the analysis and prepared the manuscript with contributions from all-authors.

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# Tables

#### Table 1: Basic configurations of participating modeling systems in MICS-Asia III

No	Horizontal resolution	Vertical resolution	First layer height	Horizontal advection	Vertical advection	Horizontal Diffusion	Vertical Diffusion	Gas phase chemistry	Aerosol processes	Dry depositiono f gases	Wet deposition of gases	Meteorology	Boundary condition	Online (Yes or No)
М1	45km	$40\sigma_p$ level	57 m	Yamo (Yamartino, 1993)	ppm (Collella and Woodward, 1984)	multiscale	ACM2 (Pleim, 2007)	SAPRC99 (Carter, 2000)	Aero6 (Binkowski and Roselle, 2003)	Wesely (1989)	Henry's law	Standard <sup>a</sup>	GEOS-Chem (Martin et al., 2002)	No
M2	45km	$40\sigma_p$ level	57 m	Yamo	ppm	multiscale	ACM2	SAPRC99	Aero6	Wesely (1989)	Henry's law	Standard <sup>a</sup>	Default	No
M3	45km	$40\sigma_p \; {\rm level}$	57 m	Yamo	Yamo	multiscale	ACM2	CB05 (Yarwood et al., 2005)	Aero5	Wesely (1989)	Henry's law	Standard <sup>a</sup>	GEOS-Chem	No
M4	45km	$40\sigma_p$ level	57 m	ppm	ppm	multiscale	ACM2_ inline	SAPRC99	Aero5	Wesely (1989)	Henry's law	Standard <sup>a</sup>	CHASER (Sudo et al., 2002a)	No
M5	45km	$40\sigma_p$ level	57 m	ppm	ppm	multiscale	ACM2_ inline	SAPRC99	Aero5	M3DRY (Pleim et al., 2001)	Henry's law	Standard <sup>a</sup>	CHASER	No
M6	45km	$40\sigma_p \; {\rm level}$	57 m	Yamo	Yamo	multiscale	ACM2_ inline	SAPRC99	Aero5	M3DRY	ACM	Standard <sup>a</sup>	CHASER	No
М7	45km	$40\sigma_p$ level	29 m	Monotonic	Monotonic	no diffusion	no diffusion	RACM-ESRL with KPP (Goliff et al.,2013)	MADE (Ackerman n et al., 1998)	Wesely (1989)	Henry's law	WRF/NCEP <sup>a</sup>	Default	No
M8	45km	$40\sigma_p$ level	57 m	5 <sup>th</sup> order Monotonic	3 <sup>th</sup> order Monotonic	MYJ	MYJ	RACM with KPP	MADE	Wesely (1989)	AQCHEM	WRF/NCEP <sup>a</sup>	CHASER	Yes
M9	45km	$40\sigma_p \; {\rm level}$	16 m	5 <sup>th</sup> order Monotonic	3 <sup>th</sup> order Monotonic	Smagorinsk y first order closure	YSU (Hong et al., 2006)	RADM2 (Stockwell et al., 1990)	MADE	Wesely (1989)	Easter et al., (2004)	WRF/NCEP <sup>a</sup>	GEOS-Chem	Yes
M10	45km	$60\sigma_p$ level	44 m	Monotonic	3 <sup>th</sup> order Monotonic	2th order Monotonic	YSU	RADM2	GOCART	Wesely (1989)	Grell	WRF/ MERRA2 <sup>a</sup>	MOZART + GOCART <sup>b</sup>	No
M11	45km	$20\sigma_z$ level	50 m	Walcek and Aleksic (1998)	Walcek and Aleksic (1998)	multicale	K-theory	CBMZ (Zaveri et al.,1999)	ISORROPI A1.7 (Nenes et al.,1998)	Wesely (1989)	Henry's law	Standard <sup>a</sup>	CHASER	No







M12	45km	40 $\sigma_p$ level	54 m	Walcek and Aleksic (1998)	Walcek and Aleksic (1998)	FTCS	FTCS	SAPRC99	Kajino et al. (2012)	Zhang et al. (2003)	Henry's law	Standard <sup>a</sup>	CHASER	No
M13	0.5°×0.667°	$47\sigma_p$ level	60 m	ppm	Huynh/Van Leer/Lin full monotonicity constraint	Lin and McElroy, 2010	Lin and McElroy, 2010	Tropchem	ISORROPI A2.0 (Fountoukis and Nenes, 2007)	Wesely	Henry's law	GEOS-5ª	Geos-Chem	No
M14	64km	$15\sigma_z$ level	100 m	ppm	ppm	multiscale	ACM2	SAPRC99	ISORROPI A1.7	Wesely (1989)	Henry's law	RAMS/NCEP <sup>a</sup>	Geos-Chem	No

<sup>a</sup> Standard represents the reference meteorological field provided by MICS-Asia III project; WRF/NCEP and WRF/MERA represents the meteorological field of the participating model itself, which was run by WRF driven by the NCEP and Modern Era Retrospective-analysis for Research and Applications (MERA) reanalysis dataset. RAMS/NCEP is the meteorology field run by RAMS driven by the NCEP reanalysis dataset. <sup>b</sup> Boundary conditions of M10 are from MOZART and GOCART (Chin et al., 2002; Horowitz et al., 2003), which provided results for gaseous pollutants and aerosols, respectively.

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Species	Dagiona	Statistics	Model													
	Regions		M1	M2	M4	M5	M6	M7	M8	M9	M10	M11	M12	M13	M14	Ense
NO		R(spatial) <sup>a</sup>	0.63	0.67	0.67	0.67	0.67	0.70	0.70	0.59	0.57	0.66	0.69	-	0.70	0.67
	NCP	R(temporal) <sup>b</sup>	0.82	0.92	0.93	0.86	0.92	0.81	0.28	0.85	0.95	0.75	0.90	-	0.96	0.91
		MBE	-4.11	-5.66	-6.54	1.86	-5.12	-5.04	3.30	8.28	-2.45	0.00	-3.81	-	-2.99	-1.86
		NMB(%)	-17.8	-24.5	-28.4	8.0	-22.2	-21.9	14.2	35.9	-10.6	0.02	-16.5	-	-13.0	-8.0
		RMSE	7.40	8.25	8.79	6.75	8.01	7.55	6.54	12.74	7.72	6.37	7.38	-	6.68	6.36
1102		R(spatial) <sup>a</sup>	0.12	0.06	0.07	0.07	0.06	0.12	0.20	0.38	0.00	0.08	0.12	-	0.02	0.10
		R(temporal) <sup>b</sup>	0.93	0.80	0.86	0.88	0.79	0.68	0.83	0.95	0.74	0.74	0.75	-	0.52	0.86
	PRD	MBE	-6.73	-9.84	-7.21	1.96	-6.66	-3.99	3.24	-7.61	-1.84	3.02	-5.49	-	-5.03	-3.85
		NMB(%)	-30.1	-44.0	-32.3	8.8	-29.8	-17.9	14.5	-34.0	-8.2	13.5	-24.6	-	-22.5	-17.2
		RMSE	11.31	13.14	12.00	10.80	11.84	10.60	8.73	10.69	10.72	10.51	11.68	-	12.00	10.15
	NCP	R(spatial) <sup>a</sup>	0.35	0.48	0.27	0.34	0.36	0.22	0.19	0.48	0.49	0.33	0.35	-0.13	0.29	0.37
		R(temporal) <sup>b</sup>	0.94	0.96	0.92	0.22	0.90	0.77	0.94	0.92	0.82	0.85	0.94	0.85	0.88	0.92
		MBE	-1.53	-1.35	-1.59	-1.69	-1.52	-1.64	-1.29	-1.16	-1.55	-1.37	-1.38	-1.53	-1.51	-1.47
		NMB(%)	-68.9	-60.9	-71.4	-76.2	-68.2	-73.7	-58.2	-52.0	-70.0	-61.6	-62.3	-68.9	-68.0	-66.2
CO		RMSE	1.71	1.54	1.77	1.86	1.70	1.82	1.51	1.36	1.74	1.57	1.58	1.74	1.70	1.66
00	PRD	R(spatial) <sup>a</sup>	0.04	-0.24	-0.25	-0.23	-0.22	-0.05	0.08	0.55	-0.02	-0.01	-0.22	0.09	-0.21	-0.06
		R(temporal) <sup>b</sup>	0.96	0.91	0.93	0.84	0.95	0.90	0.90	0.96	0.83	0.87	0.93	0.76	0.82	0.94
		MBE	-0.66	-0.64	-0.65	-0.64	-0.62	-0.64	-0.51	-0.57	-0.50	-0.51	-0.58	-0.52	-0.67	-0.59
		NMB(%)	-68.4	-67.0	-67.0	-66.7	-64.7	-66.5	-53.3	-59.7	-52.3	-52.7	-60.7	-54.1	-69.6	-61.7
		RMSE	0.70	0.70	0.70	0.69	0.67	0.69	0.57	0.62	0.56	0.57	0.64	0.58	0.72	0.65
NH3	NCP	R(spatial) <sup>a</sup>	0.72	0.70	0.69	0.70	0.71	0.65	-	0.70	0.57	0.62	0.67	0.61	0.58	0.69
		R(temporal) <sup>b</sup>	-0.48	-0.22	-0.45	-0.55	-0.41	0.04	-	-0.19	0.64	0.08	-0.37	0.65	-0.04	-0.17
		MBE	-0.69	2.95	-6.14	-6.61	-3.89	4.94	-	21.8	10.5	-0.07	0.31	-5.19	-12.2	0.47
		NMB(%)	-3.8	16.1	-33.5	-36.0	-21.2	26.9	-	118.7	57.1	-0.4	1.69	-28.3	-66.3	2.59
		RMSE	7.20	10.04	8.95	9.24	7.48	8.78	-	29.24	13.48	8.30	7.33	8.82	14.48	7.20

### Table 2: Statistics of simulated annual mean concentrations sampled from different sites in NCP and PRD.

<sup>a</sup> R(spatial) represents the spatial correlation coefficients between simulated and observed concentrations sampled from different stations in NCP or PRD;

<sup>b</sup> R(temporal) represents the temporal correlation coefficients between simulated and observed monthly mean concentrations from January to December in 2010;





# Figures



Figure 1: Modeling domains of the participated models except M13 and M14 along with spatial distributions of the total emissions of (a) NO<sub>x</sub>, (b) CO and (c) NH<sub>3</sub> in 2010 provided by MICS-Asia III (upper panel), and the distributions of observation stations of (d) NO<sub>2</sub> and CO over NCP and PRD, as well as (e) NH<sub>3</sub> over NCP (lower panel). The horizontal resolution is 45km×45km. Note that domains of M13 and M14 are shown in fig. 7 and only six of nineteen observational sites (blue) over NCP have CO measurements.







Figure 2: Boxplot of simulated and observed annual mean NO<sub>2</sub>, CO and NH<sub>3</sub> concentrations sampled from different stations over NCP (a, c, e) and PRD (b, d). The outlier was defined as values larger than  $q_3 + 1.5 \times (q_3 - q_1)$  or less than  $q_1 - 1.5 \times (q_3 - q_1)$ , where  $q_3$  denotes the 75th percentile, and  $q_1$  the 25th percentile. This approximately corresponds to 99.3 percent coverage if the data are normally distributed.







Figure 3: Timeseries of regional mean NO<sub>2</sub>, CO concentrations over NCP (a, c) and PRD (b, d) as well as NH<sub>3</sub> concentrations over NCP (e) from January to December in year 2010.







Figure 4: Timeseries of NH<sub>3</sub> emissions over NCP provided by MICS-Asia III on a horizontal resolution of 45km×45km from January to December in year 2010.

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Figure 5: Timeseries of the multi-model simulated total ammonium ( $NH_x = NH_3 + NH_4^+$ ) in atmosphere along with the ratio of gaseous NH<sub>3</sub> to the total ammonium over NCP from January to December in year 2010.







Figure 6: Spatial distribution of the annual mean NO<sub>2</sub> concentrations from each modeling results of MICS-Asia III. Note that M13 are not included in this figure.







Figure 7: Spatial distribution of the annual mean CO concentrations from each modeling results of MICS-Asia III.







Figure 8: Spatial distribution of the annual mean NH<sub>3</sub> concentrations from each modeling results of MICS-Asia III.







Figure 9: Spatial distribution of the standard deviation of (a) NH<sub>3</sub>, (b) CO and (c) NO<sub>2</sub> multi-model predictions in MICS-Asia III, as well as the corresponding distribution of CV on the annual and seasonal basis.