A potential source of atmospheric sulfate from O₂⁻-induced SO₂ oxidation by ozone

Narcisse Tchinda Tsona and Lin Du

Environment Research Institute, Shandong University, Binhai Road 72, 266237 Shandong, China

5   *Correspondence to*: Lin Du (lindu@sdu.edu.cn)

Abstract. It was formerly demonstrated that O_2SOO^- forms at collisions rate in the gas-phase as a result of SO_2 reaction with O_2^-. Hereby, we present a theoretical investigation of the chemical fate of O_2SOO^- by reaction with O_3 in the gas-phase, based on *ab initio* calculations. Two main mechanisms were found for the title reaction, with fundamentally different products: (i) formation of a van der Waals complex followed by electron transfer and further decomposition to $O_2 + SO_2 + O_3^-$ and (ii)

10   formation of a molecular complex from O_2 switching by O_3, followed by SO_2 oxidation to SO_3^- within the complex. Both reactions are exergonic, but separated by relatively low energy barriers. The products in the former mechanism would likely initiate other SO_2 oxidations as shown in previous studies, whereas the latter mechanism closes a path wherein SO_2 is oxidized to SO_3^-. The latter reaction is atmospherically relevant since it forms the SO_3^- ion, hereby closing the SO_2 oxidation path initiated by O_2^-. The main atmospheric fate of SO_3^- is nothing but sulfate formation. Exploration of the reactions kinetics

15   indicates that the path of reaction (ii) is highly facilitated by humidity. For this path, we found an overall rate constant of 4.0×10^{-11} cm³ molecule⁻¹ s⁻¹ at 298 K and 50% relative humidity. The title reaction provides a new mechanism for sulfate formation from ion-induced SO_2 oxidation in the gas-phase and highlights the importance of including such mechanism in modelling sulfate-based aerosol formation rates.

1 Introduction

The chemistry of sulfur is highly important in the atmosphere. Through its oxidation products, sulfur participates in the formation of secondary atmospheric aerosols, clouds and acid rain. Sulfur dioxide (SO_2), the most abundant sulfur-containing molecule in the atmosphere, is known to oxidize in the gas-phase following different mechanisms to form sulfuric acid (H_2SO_4) as the final oxidation species. The main SO_2 oxidizers include the hydroxyl radical (OH) (Seinfeld and Pandis, 2016), stabilized Criegee intermediates (Welz et al., 2012; Mauldin III et al., 2012; Vereecken et al., 2012), mineral dust (Harris et al., 2013) and atmospheric ions (Fehsenfeld and Ferguson, 1974; Enghoff et al., 2012; Tsona et al., 2015). Mechanisms involving OH and Criegee intermediates lead to the formation of SO_3 that ultimately forms H_2SO_4 (Larson et al., 2000), whereas SO_2 reactions with ions are generally more complex. H_2SO_4 is known to be the main driving species in atmospheric aerosols formation and its formation is critical in the determination of aerosol formation rates (Nieminen et al., 2009; Sipila et al., 2010), yet the role of ions in this process has been well established (Yu, 2006; Yu and Turco, 2000, 2001; Enghoff and Svensmark, 2008; Kirkby et al., 2011; Wagner et al., 2017; Yan et al., 2018).

The immediate products of SO_2 oxidation by ions are mostly sulfur oxides ions intermediates (Fehsenfeld and Ferguson, 1974; Möhler et al., 1992; Bork et al., 2012; Tsona et al., 2014) that are susceptible of triggering new reactions or recombining with oppositely charged counterparts to form neutral species. Some of these ions, namely SO_3^-, SO_4^-, and SO_5^-, were detected at relatively high concentrations in the ambient atmosphere (Ehn et al., 2010) and in chamber experiments of SO_2 ionic oxidation studies (Nagato et al., 2005; Hvelplund et al., 2013; Kirkby et al., 2011; Kirkby et al., 2016). The chemical fate of most sulfur oxides anions is relatively known. Bork et al. showed that SO_3^- can form SO_3, the precursor for H_2SO_4, through electron transfer to ozone (O_3) (Bork et al., 2012). SO_3^- can equally react with O_2 to form SO_5^- whose atmospheric outcome by reaction with O_3 is H_2SO_4 formation (Bork et al., 2013). It was also speculated from chamber studies that SO_5^- could form and stabilize clusters with sulfuric acid in the gas-phase (Kirkby et al., 2011). Reliable predictions of the outcomes of these ions require an exact knowledge of their chemical structures since interactions between molecules or ions depend both on their physical and chemical properties. A previous study demonstrated that two forms of SO_4^- separated by a high energy barrier may exist in the atmosphere (Tsona et al., 2014): the sulfate radical ion henceforth indicated as SO_4^-, and the peroxy form, O_2SOO^-. Formerly, the two isomers were often misleadingly attributed exclusively to the sulfate radical ion, the most stable form of SO_4^-. However, their reactive properties greatly differ (Fehsenfeld and Ferguson, 1974).

The formation mechanisms of SO_4^- in the gas-phase have been largely unknown but, recent studies showed that this ion can be formed by SO_5^- reaction with O_3 (Bork et al., 2013) and in a O_2SOO^- isomerization process catalyzed by NO (Tsona et al., 2018). SO_4^- can also be produced during the chemical transformation of organic compounds, triggered by sulfate salts (Noziere et al., 2010), whereas O_2SOO^- is formed at collision rates upon SO_2 reaction with O_2^- (Fahey et al., 1982; Tsona et al., 2014).

The sulfate radical ion is believed to react with unsaturated compounds to form organosulfates, a major component of secondary organic aerosol (Surratt et al., 2007;Surratt et al., 2008;Schindelka et al., 2013). Using first-principles calculations, it was demonstrated that SO_4^- acts as a catalyst in SO_2 oxidation to SO_3 by O_3 in the gas-phase and hence, plays a role in atmospheric aerosol formation (Tsona et al., 2015;Tsona et al., 2016). The chemistry of O_2SOO^- is largely unknown but, is

5 potentially important for sulfur chemistry and atmospheric aerosol formation. Fehsenfeld and Ferguson found that O_2SOO^- can be decomposed by NO_2 into NO_3^- and SO_3 (Fehsenfeld and Ferguson, 1974) and it was formerly demonstrated that in the presence of nitrogen oxides (NOx = NO_2 + NO), O_2SOO^- can be converted into sulfates (Tsona et al., 2018). In mildly polluted environments, the concentration of O_3 can be few orders of magnitude higher than that of NOx and the chemical fate of O_2SOO^- would then also greatly depend on collisions with O_3. in such environments, O_2SOO^- could experience much more collisions

10 with O_3 than with NOx.

Hereby, we investigate the reaction between $O_2SOO^-\cdots(H_2O)_{0-1}$ and O_3 using *ab initio* calculations. By determining the reactions thermodynamics and kinetics, we examine the possible pathways for the reaction and propose the most probable outcome of $O_2SOO^-\cdots(H_2O)_{0-1}$ based on O_3 reaction. Implications of the most relevant pathways in aerosol formation are

15 discussed.

2 Methods

2.1 Geometry optimizations, thermochemical and charge analysis

All molecules and stationary points in the energy surface were optimized using density functional theory (DFT) based on the M06-2X density functional (Zhao and Truhlar, 2008), and the aug-cc-pVTZ basis set (Dunning Jr et al., 2001). This functional

20 has successfully proven to be adequate for reactions involving transition state (TS) configurations (Elm et al., 2012, 2013b, a). Harmonic vibrational frequencies analysis on the optimized structures were performed (at 298 K and 1 atm) using the M06-2X/aug-cc-pVTZ method under the harmonic oscillator-rigid rotor approximation. These calculations ensured that the obtained stationary points were minima or TS and, also, provided the thermal contributions to the Gibbs free energy and the enthalpy.

25 Transition states structures were initially located by scanning the reactants configurations. The best TS guesses out of the scans were then refined using the synchronous transit quasi Newton method (Peng et al., 1996), and the final TS structures underwent internal reaction coordinate calculations (Fukui, 1981) to ensure they connected the reactants to desired products.

The electronic energies of the M06-2X/aug-cc-pVTZ optimized geometries were corrected with the CCSD(T) method (Purvis

30 and Bartlett, 1982) in conjunction with the aug-cc-pVTZ basis set. The Gibbs free energies, *G*, of all relevant species were then calculated as

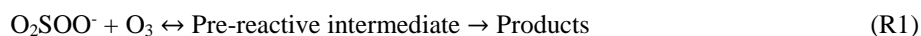
$$G = E_{CCSD(T)} + G_{therm} \tag{1}$$

where $E_{CCSD(T)}$ is the electronic energy calculated with the CCSD(T)/aug-cc-pVTZ method and G_{therm} is the thermal contribution to the Gibbs free energy, calculated at the M06-2X/aug-cc-pVTZ level of theory. All geometry optimizations, harmonic vibrational frequencies analysis and electronic energies correction calculations were carried out in the Gaussian 09 package (http://gaussian.com/).

To analyse the distribution of the excess electronic charge over different species and fragments in the optimized systems, we used the Atoms-in-Molecules charge partitioning method presented by Bader (Bader, 1998). As input, this method requires electronic density and nuclear coordinates from electronic structure calculations. We used the approach implemented in the algorithm developed by Henkelman and co-workers, which has been shown to work well both for charged and water-containing systems (Tang et al., 2009; Bork et al., 2011; Henkelman et al., 2006).

2.2 Reactions kinetics

Regardless of the presence of water, the reaction between O_2SOO^- and O_3 begins by forming different van der Waals pre-reactive intermediates, depending on the orientation of the reactants at impact. The pre-reactive intermediate could either decompose to different species or react further through a transition state configuration to form new products:

$$O_2SOO^- + O_3 \leftrightarrow \text{Pre-reactive intermediate} \rightarrow \text{Products} \tag{R1}$$

The traditional approach to determine the rate constant of reaction (R1) relies on the steady-state approximation and leads to the following equation:

$$k = k_{coll} \frac{k_{reac}}{k_{reac} + k_{evap}} \tag{2}$$

where k_{coll} is the collision frequency for O_2SOO^--O_3 collisions, k_{evap} is the rate constant for the evaporation of the pre-reactive intermediate back to initial reactants, and k_{reac} is the unimolecular rate constant for the reaction of the pre-reactive intermediate to the products. Moreover, assuming that $k_{evap} \gg k_{reac}$, the rate constant of reaction (R1) becomes $k = K_{eq}k_{reac}$ over a range of temperatures, with K_{eq} being the equilibrium constant of formation of the pre-reactive intermediate from the reactants. This consideration is, however, not valid for barrierless reactions, since the pre-reactive intermediate seldom thermally equilibrates. For such reactions, a two-transition state theory has been introduced, treating two distinct transition state bottlenecks that define the net reactive flux (Klippenstein et al., 1988; Georgievskii and Klippenstein, 2005; Greenwald et al., 2005). The first bottleneck, the "outer" transition state, occurs in the association of the initial reactants to form the pre-reactive intermediate, whereas the second bottleneck, the "inner" transition state, occurs in the transformation of the pre-reactive intermediate to the

products. Based on this theory, the overall rate constant (k) for a reaction channel is expressed in terms of the outer (k_{out}) and inner (k_{in}) rate constants as:

$$\frac{1}{k} = \frac{1}{k_{out}} + \frac{1}{k_{in}} \tag{3}$$

The outer transition state is treated by the long-range transition state theory approach (Georgievskii and Klippenstein, 2005), while the inner transition state is resolved by the transition state theory. It has been demonstrated that for interactions between ions and neutral molecules, due to their long-range attraction, the collision cross section is larger than would be measured from the physical dimensions of the colliding species (Kupiainen-Määttä et al., 2013). To account for this phenomenon, the outer

10 rate constant was determined from the ion-dipole parametrization of Su and Chesnavich who performed trajectory simulations of collisions between a point charge and a rigidly rotating molecule (Su and Chesnavich, 1982). This parametrization is equivalent to a Langevin capture rate constant (k_L) scaled by a temperature-dependent term and was found to provide good agreement with experiments (Kupiainen-Määttä et al., 2013). It is given as

$$k_{out} = k_L \times \left(\frac{(x+0.5090)^2}{10.526} + 0.9754 \right) \tag{4}$$

where $k_L = q\mu^{-1/2}(\pi\alpha/\varepsilon_0)^{1/2}$, $x = \mu_D/(8\pi\varepsilon_0\alpha k_B T)^{1/2}$, q is the charge of the ion, μ is the reduced mass of the colliding species, ε_0 is the vacuum permittivity, α and μ_D are the polarizability and dipole moment of the neutral molecule, k_B is Boltzmann's constant, and T is the absolute temperature. The inner rate constant can be written as:

$$k_{in} = \frac{k_B T}{hc^0} \times \exp\left(-\frac{\Delta G^{\#}}{RT}\right) \tag{5}$$

where $\Delta G^{\#}$ is the Gibbs free energy barrier separating the pre-reactive intermediate and the products, h is the Planck's constant, R is the molar gas constant, and c^0 is the standard gas-phase concentration.

25

3 Results and discussion

Starting with optimized structures of $O_2SOO^{-}\cdots(H_2O)_{0-1}$ and O_3 shown in Fig. S1, a series of geometry optimizations were performed on $O_2SOO^{-}\cdots(H_2O)_{0-1} + O_3$ interaction, taking into account different spatial orientations of the reactants at impact. These optimizations led to two main chemical processes, depending on the initial orientation of the reactants, with potentially

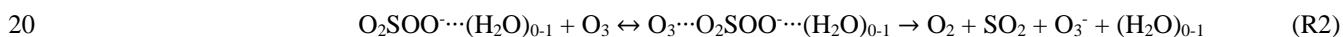
30 different outcomes. The first process is the formation of a van der Waals complex followed by its direct decomposition to other

species. The second process is the barrierless formation of a molecular complex in which the SO_2 entity of $O_2SOO^-\cdots(H_2O)_{0-1}$ is oxidized to SO_3^-.

3.1 Cluster formation and decomposition of $O_2SOO^-\cdots(H_2O)_{0-1}$

As O_3 approaches $O_2SOO^-\cdots(H_2O)_{0-1}$ towards the oxygen atoms of the peroxy fragment or the oxygen atom of the SO_2 entity, the immediate outcome of $O_2SOO^-\cdots(H_2O)_{0-1}$ and O_3 collisions is the formation of the van der Waals $O_3\cdots O_2SOO^-\cdots(H_2O)_{0-1}$ complex in which O_3 interacts with O_2SOO^-. Among the different stable configurations found upon optimizations, we solely report the most stable one with respect to the Gibbs free energy, which is henceforth denoted RC1 and RCW1 for the unhydrated and monohydrated, respectively, shown in Fig. 1. Exploration of RC1 and RCW1 structures reveals that $O_2SOO^-\cdots(H_2O)_{0-1}$ basically keeps its configuration upon clustering with O_3. Their formations are endergonic at most atmospheric temperature, with Gibbs free energy changes of 4. 5 and 4.7 kcal mol^{-1}, respectively, at 298 K. These values indicate that, when formed, these complexes would not live long and will rather decompose either to initial reactants or to different species. Inspecting the vibrational modes of RC1 and RCW1, two vibrations are found that would clearly lead to the dissociation of O_2SOO^- within the cluster. The analysis of the charge distribution over $O_3\cdots O_2SOO^-\cdots(H_2O)_{0-1}$ shows that the extra electron initially located on O_2SOO in the reactants has partially migrated to the O_3 molecule in the van der Waals product complex, as can be observed in Fig. S2. This is as expected, given the high electronegativity of O_3 relative to those of O_2 and SO_2 (Rothe et al., 1975). The charge distribution over the different atoms of the optimized complex is weakly affected by the presence of water, as previously demonstrated by Bork and co-workers (Bork et al., 2011).

The most likely fates of RC1 and RCW1 are, therefore, decomposition into O_2, SO_2 and O_3^- as follows:

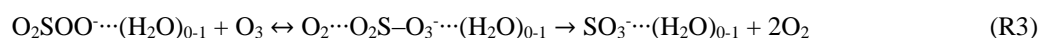
$$O_2SOO^-\cdots(H_2O)_{0-1} + O_3 \leftrightarrow O_3\cdots O_2SOO^-\cdots(H_2O)_{0-1} \rightarrow O_2 + SO_2 + O_3^- + (H_2O)_{0-1} \tag{R2}$$

The numerical values of the formation energies of all intermediate species in reaction (R2) are given in Table 1 and the energy surfaces are plotted in Fig. 2. RC1 and RCW1 decompositions are highly exergonic at 298 K, occurring with -18.1 and -16.7 kcal mol^{-1} Gibbs free energy changes, respectively. These processes are, therefore, likely to occur in the atmosphere upon formation of $O_3\cdots O_2SOO^-\cdots(H_2O)_{0-1}$.

The limiting step in reaction (R2) is the formation of RC1 and RCW1, whose formation energies indicated above can then be considered as the only barrier to the formation of $O_2 + SO_2 + O_3^-$. This leads to overall rate constants (according to Eq. 5) of 1.4×10^{-10} and 9.9×10^{-11} cm^3 molecule^{-1} s^{-1} at 298 K for the unhydrated and monohydrated reaction, respectively. Both reactions are, in principle, collision-limited and the effect of hydration on the kinetics is found to be negligible. The atmospheric relevance of reaction (R2) has been determined earlier (Bork et al., 2012; Bork et al., 2013; Enghoff et al., 2012).

3.2 $O_2SOO^-\cdots(H_2O)_{0-1}$ reaction with O_3

When O_3 approaches $O_2SOO^-\cdots(H_2O)_{0-1}$ from the sulfur atom side, the formation of a more stable cluster than found above prevails. The incoming O_3 molecule strongly interacts with $O_2SOO^-\cdots(H_2O)_{0-1}$ by forming a coordination bond with the sulfur atom and hereby, inducing the ejection of the O_2 molecule that remains in interaction with the remainder of the system. This

5   process leads to the formation of the $O_2\cdots O_2S–O_3^-\cdots(H_2O)_{0-1}$ complex which further transforms, through an intramolecular SO_2 oxidation, into $SO_3^-\cdots(H_2O)_{0-1} + 2O_2$ according to the following equation:

$$O_2SOO^-\cdots(H_2O)_{0-1} + O_3 \leftrightarrow O_2\cdots O_2S–O_3^-\cdots(H_2O)_{0-1} \rightarrow SO_3^-\cdots(H_2O)_{0-1} + 2O_2 \qquad (R3)$$

10   The configurations of the most stable intermediate structures in reaction (R3) are given in Fig. 1. The charge analysis indicated that the oxygen molecules released in the products are in the singlet state. Though the necessity to determine the electronic structure of the O_2 molecule in the singlet sate ($^1\Delta_g$) has been demonstrated to be useful (Buttar and Hirst, 1994), obtaining a reliable electronic energy for $O_2(^1\Delta_g)$ is difficult (Drougas and Kosmas, 2005). An alternative approach to determine this energy is to add the experimental energy spacing (22.5 kcal mol^{-1}) between triplet and singlet states of O_2 to the computed

15   electronic energy of the triplet O_2 (Schweitzer and Schmidt, 2003; Drougas and Kosmas, 2005). We used this approach to determine the energies of the products of reaction (R3). The numerical values of the formation energies of all intermediate species in reaction (R3) are thus given in Table 1 and the energy surfaces are plotted in Fig. 2. The most stable optimized structures of $O_2\cdots O_2S–O_3^-\cdots(H_2O)_{0-1}$ according to our calculations are denoted as RC2 and RCW2 for the unhydrated and monohydrated systems, respectively and are shown in Fig. 1. Regardless of the presence of water, the $O_2\cdots O_2SO_3^-$ configuration

20   results from O_2 being switched by O_3 in the O_2SOO^- molecular ion. In the optimized $O_2\cdots O_2S–O_3^-$ structure, the O3 atom of O_3 points towards the S atom of O_2SOO^-, forming S–O3 bonds at distances of 1.90 and 1.87 Å in the absence and presence of water, respectively. These bonds are coordination bonds in nature since the S-O covalent bond in e.g., 1.43 Å in SO_3 and 1.46 Å in H_2SO_4. The S-O3 coordination bond distances in RC2 and RCW2 are shorter by 0.04 and 0.03 Å than $O_2S–OO^-$ bond distances in unhydrated and monohydrated O_2SOO^- forms. This indicates stronger interaction between O_3 and SO_2, and hence

25   higher stability of $O_2\cdots O_2S–O_3^-$ relative to O_2SOO^-.

The formations of RC2 and RCW2 are highly exergonic, with Gibbs free energy changes at 298 K of -14.7 and -12.4 kcal mol^{-1}, respectively. These values, with corresponding electronic energies and enthalpies are shown in Table 1. These Gibbs free energy changes for the formation of RC2 and RCW2 are about 19 kcal mol^{-1} lower than those of RC1 and RCW1 at similar

30   conditions, indicating the higher stability of RC2 and RCW2, and the highly favourable switching reaction at ambient conditions. SO_2 oxidation can readily occur within the $O_2\cdots O_2S–O_3^-\cdots(H_2O)_{0-1}$ cluster and lead to SO_3^- formation. In principle, to form the products of reaction (R3), the O3 atom of the O_3 fragment in $O_2\cdots O_2S–O_3^-\cdots(H_2O)_{0-1}$ transfers, through transition state configurations, to SO_2 and form SO_3^- followed by the ejection of the O_2 molecule. The transition states are denoted TS2

and TSW2 for the unhydrated and monohydrated systems, respectively, and their structures are presented in Fig. 1. While RC2 and RCW2 are formed with similar Gibbs free energies within 2 kcal mol^{-1} difference, the formation Gibbs free energies of their transition states at similar conditions greatly differ. TS2 is located at 10 kcal mol^{-1} Gibbs free energy above RC2, while TSW2 is located at -4 kcal mol^{-1} below RCW2. It is speculated that the low energy barrier in the monohydrated reaction is the

5   result of a strong stabilisation of the transition state due to hydration. The S-O3 bonds in RCW2 and TSW2 are shorter by ~0.03 Å than in RC2 and TS2.

Based on TS2 and TSW2 energies, the unimolecular decomposition of $O_2\cdots O_2S–O_3^-$ at 298 K was found to occur at rate constants of 3.1×10^5 s^{-1} and 6.5×10^{15} s^{-1}, corresponding to atmospheric lifetimes of 3.3×10^3 ns and 1.5×10^{-7} ns for the

10   unhydrated and monohydrated systems, respectively. The obtained short lifetimes indicate that $O_2\cdots O_2S–O_3^-$ would not live long enough to experience collisions with other atmospheric oxidants. It should be noted that few to no collisions with nitrogen can, however, be achieved. It follows that the most likely outcome of $O_2\cdots O_2S–O_3^-$ is decomposition to the products of reaction (R3), which are formed with about -23 kcal mol^{-1} overall Gibbs free energy at 298 K. The net reaction is an O_2^--initiated SO_2 oxidation to SO_3^- by O_3.

The charge analysis clearly indicates that the extra electron, initially confined on the atoms of the O_3 fragment of $O_2\cdots O_2S–O_3$ (RC2 structure), is partially distributed over other atoms in the transition state (TS2) and is finally located on SO_3 in the product (see **Fig. 3**). Apparent also from **Fig. 3** is the antibonding orbital of the singlet O_2 molecule.

20   The overall rate constants of reaction (R3), determined at 298.15 K using Eq. (3), are 1.3×10^{-14} and 8.0×10^{-10} cm^3 molecule^{-1} s^{-1} for the unhydrated and monohydrated reactions, respectively. The values of the different components (k_{out} and k_{in}) are listed in Table S1 of the Supplement. It is observed from Table S1 that the inner transition sate provides the dominant bottleneck to the rate constant of the unhydrated reaction, whereas the outer transition state provides the dominated bottleneck to the rate constant of the monohydrated reaction.

25

The effective effect of water on the rate constant can be evaluated by taking into account the stability of $O_2SOO^-\cdots H_2O$ (which is formed at the entrance channel of the reaction in the presence of water before colliding with O_3) and the equilibrium vapor pressure of water. Starting from the definition of the reaction rate for the hydrated reaction,

$$J_{(R3w)} = k_{(R3w)} \times [O_2SOO^-\cdots H_2O] \times [O_3] \qquad (6)$$

$$= k_{(R3w)}^{eff} \times [O_2SOO^-]\times[O_3] \qquad (7)$$

where $k_{(R3w)}$ is the overall rate constant for the hydrated reaction, determined using Eq. (3), $k_{(R3w)}^{eff}$ is the effective reaction rate constant calculated as $k_{(R3w)}^{eff} = k_{(R3w)} \times K_{eq} \times p_{H2O}$. Keq is the equilibrium constant for the $O_2SOO^- + H_2O \leftrightarrow O_2SOO^{-}\cdots H_2O$ reaction and $[H_2O]$ the actual water vapor pressure. Details on the determination of K_{eq} and p_{H2O} are given in the Supplement.

5   At 298 K and 50 % relative humidity, the effective rate constant of the monohydrated reaction is 1.7×10^{-10} cm^3 molecule^{-1} s^{-1}, four orders of magnitude higher than the rate constant of the unhydrated reaction. Therefore, water plays a catalytic role on the kinetics of reaction (R3). The net rate constant of reaction (R3) can be obtained by weighing the rate constants of the unhydrated and monohydrated reactions to corresponding equilibrium concentrations of $O_2\cdots O_2S–O_3^-$ hydrates. Using the law of mass action, we find that $O_2\cdots O_2S–O_3^-$ mostly exists as a dry species, constituting 77% of the total population, whereas the

10   monohydrated species is formed at 23%. The net rate constant of reaction (R3) is then determined to be 4.0×10^{-11} cm^3 molecule^{-1} s^{-1} at 298 K.

Considering only the unhydrated process of reaction (R3), the rate constant is 4-5 orders of magnitude lower than the rate constant obtained for the $SO_2 + O_3^- \rightarrow SO_3^- + O_2$ reaction (Fehsenfeld and Ferguson, 1974; Bork et al., 2012). Despite this

15   difference, the oxidation process follows a similar mechanism to the one presented by Bork et al. for the $SO_2 + O_3^- \rightarrow SO_3^- + O_2$ reaction, consisting of the oxygen transfer from O_3 to SO_2 (Bork et al., 2012). The discrepancy between the two results is associated with the effect of the presence of the O–O fragment initially coordinated to SO_2 in the current study, which tends to stabilize the $O_2\cdots O_2S–O_3^-$ pre-reactive complex. The presence of the O-O fragment seemingly deactivates SO_2 for the upcoming O transfer from O_3 to form SO_3^-. However, this situation is rapidly reversed with the presence of water as the reaction

20   becomes much faster, proceeding nearly at collision rate.

3.3 Further chemistry

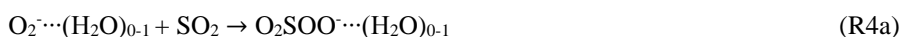
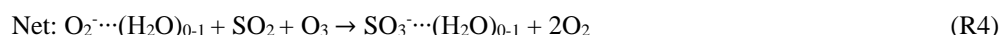
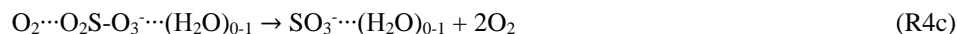
In real atmospheric and ionized conditions, despite O_2 has lower electron affinity than O_3, it would likely ionize faster than O_3 owing to its much higher concentration. Considering for example chamber experiments, upon interaction of ionizing particles

25   with the gas, electrons can transfer from one species to another and, e.g., O_2^- can form and rapidly hydrate within one nanosecond (Svensmark et al., 2007; Fahey et al., 1982). Furthermore, Fahey et al. showed that $O_2^-\cdots(H_2O)_{0-1}$ association reaction with SO_2 is faster than the electron transfer from $O_2^-\cdots(H_2O)_{0-1}$ to O_3 (Fahey et al., 1982). This means that in an ionized environment containing O_2, O_3, and SO_2, the formation of O_2SOO^- resulting from SO_2 and O_2^- association will happen faster than O_3^- formation. O_2SOO^- would react thereafter with O_3 and the following stepwise process could take place

30

$$O_2^-\cdots(H_2O)_{0-1} + SO_2 \rightarrow O_2SOO^-\cdots(H_2O)_{0-1} \tag{R4a}$$
$$O_2SOO^-\cdots(H_2O)_{0-1} + O_3 \rightarrow O_2\cdots O_2SO_3^-\cdots(H_2O)_{0-1} \tag{R4b}$$

$$O_2\cdots O_2S\text{-}O_3\cdots(H_2O)_{0\text{-}1} \rightarrow SO_3^-\cdots(H_2O)_{0\text{-}1} + 2O_2 \tag{R4c}$$

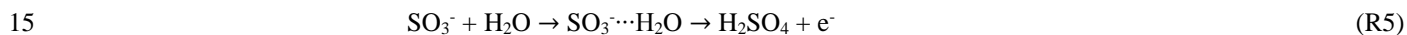
$$\text{Net: } O_2^-\cdots(H_2O)_{0\text{-}1} + SO_2 + O_3 \rightarrow SO_3^-\cdots(H_2O)_{0\text{-}1} + 2O_2 \tag{R4}$$

The Gibbs free energy change of this net reaction at 298 K is about -20 kcal mol^{-1} more negative than that of the $SO_2 + O_3^- \rightarrow$ SO$_3^- + O_2$ reaction at similar conditions. Given that the intermediate steps of reaction (R4) are significantly fast, this reaction is believed to be an important process in most environments of SO_2 ion-induced oxidation to SO_3^- or more oxidized species. The limiting step in the process of reaction (R4) is reaction (R4c) for which an energy barrier has to be overcome before the products are released.

SO_3^- is an identified stable ion detected in the atmosphere and in experiments (Ehn et al., 2010; Kirkby et al., 2011; Kirkby et al., 2016). The chemical fate of SO_3^- is fundamentally different from that of SO_3 that forms H_2SO_4 by hydration. Likely outcomes of SO_3^- are hydrolysis, electron transfer by collision with O_3, reaction with O_2 and H_2O and, possibly, radicals, according to the following equations

$$SO_3^- + H_2O \rightarrow SO_3^-\cdots H_2O \rightarrow H_2SO_4 + e^- \tag{R5}$$

$$SO_3^- + O_3 \rightarrow SO_3 + O_3^- \tag{R6}$$

$$SO_3^- + O_2 + H_2O \rightarrow HSO_4^- + HO_2 \tag{R7}$$

$$SO_3^- + OH \rightarrow HSO_4^- \tag{R8}$$

Fehsenfeld and Ferguson showed that H_2SO_4 formation could occur in the $SO_3^-\cdots H_2O$ cluster, releasing a free electron (Reaction (R5)) (Fehsenfeld and Ferguson, 1974). Owing to the high electron affinity of O_3 relative to SO_3 (Rothe et al., 1975), the electron can transfer from SO_3 to O_3 and lead to the formation of SO_3, the precursor for sulfuric acid in the atmosphere. Moreover, the free electron released and the O_3^- formed in reactions (R5) and (R6), respectively, are potential triggers of new SO_2 oxidations with implication in aerosol formation (Svensmark et al., 2007; Enghoff and Svensmark, 2008; Bork et al., 2013). Reactions (R7) and (R8) are potential outcomes for SO_3^- as well, forming the highly stable HSO_4^- species that would terminate the oxidation process of SO_2 initiated by a free electron. Reactions (R5)–(R8) are likely competitive processes upon SO_3^- formation in the gas-phase, and their different rates would determine the number of SO_2 oxidations induced by a free electron. However, they have no other fate than HSO_4^- or H_2SO_4, the most oxidized forms of sulfur in the atmosphere, which both share many properties and play a central role in atmospheric particle formation.

Experimental studies have shown that in atmospheres heavily enriched in SO_2 and O_3, a free electron could initiate SO_2 oxidation and induce the formation of ~10^7 cm^{-3} sulfates in the absence of UV light, clearly indicating the importance of other ionic SO_2 oxidation mechanisms than UV-induced (Enghoff and Svensmark, 2008). To evaluate the importance of the mechanism presented in this study in the formation of sulfates, it is necessary to identify the scavengers that terminate the SO_2 oxidation initiated by O_2^-. Possible scavengers include radicals, NOx, acids, cations and other particles. The main ones are

likely NOx and OH, which lead to the formation of the stable NO_3^- and HSO_4^- species. If the ion concentration was known, the contribution of reaction (R4) to H_2SO_4 formation could be determined by comparing its formation rates from ionic and electrically neutral mechanisms. Alternatively, it can be assumed that reaction (R4) is terminated when the ion cluster hits a scavenger. The free electron which acts as catalyst is then scavenged. The average catalytic turnover number (TON) is defined

5   as (Kozuch and Martin, 2012):

$$TON = \frac{\text{concentration of limiting reacted molecules}}{\text{concentration of the catalyst}} \tag{8}$$

The concentration of the catalyst can be approximated to the concentration of the scavengers and, assuming $[O_3] \gg [SO_2]$, SO_2

10   is the limiting species in reaction (R4). Equation (8) can then be re-written as

$$TON \approx \frac{[SO_2]}{[OH] + [NO_x]} \tag{9}$$

The catalytic efficiency of SO_2 ion-induced oxidation is then given as

$$J_{ion} = k_{ion} \times TON \tag{10}$$

Where k_{ion} is the ion production rate. Depending on the tropospheric temperature and altitude, measurements at the CLOUD chamber experiments found $k_{ion} = 2$–100 cm^{-3} s^{-1}, covering the typical ionization range in the troposphere (Franchin et al.,

20   2015). Assuming nearly pristine conditions with $[SO_2] = 5$ ppb $= 1.2 \times 10^{11}$ molecule cm^{-3}, $[NOx] = 200$ ppt $= 4.9 \times 10^9$ molecule cm^{-3}, and $[OH] = 5.0 \times 10^5$ molecule cm^{-3} (day and night average), we determine J_{ion} in the range 5.0×10^1–2.5×10^3 cm^{-3} s^{-1}. The rate of the UV-induced SO_2 oxidation by OH is

$$J_{UV} = k_{UV} \times [SO_2] \times [OH] \tag{11}$$

25

With $k_{UV} = 1.3 \times 10^{-12}$ cm^3 molecule^{-1} s^{-1} (Atkinson et al., 2004), $J_{UV} = 7.9 \times 10^4$ molecule cm^{-3} s^{-1}, and the proportion of H_2SO_4 formed from ion-induced oxidation can be estimated from the following equation

$$\frac{[H_2SO_4]_{ion}}{[H_2SO_4]_{total}} = \frac{J_{ion}}{J_{UV} + J_{ion}} \tag{12}$$

30

We find that the contribution of ion-induced SO_2 oxidation to H_2SO_4 formation can range from 0.1 to 3.1% of the total formation rate. This estimate could be improved by considering also the SO_2 oxidation by Criegee Intermediates, another important channel for H_2SO_4 formation.

5 **3 Conclusions**

This study highlights the role of the superoxide ions (O_2^-) in SO_2 oxidation. Our previous study demonstrated that SO_2 interacts with O_2^- and forms O_2SOO^- whose atmospheric fate remains unelucidated (Tsona et al., 2014). In this study, we used ab initio calculations to assess the chemical fate of O_2SOO^- by collisions with O_3. Regardless of the presence of water, two main mechanisms are observed, leading to fundamentally different products. The first mechanism is characterized by electron

10 transfer followed by O_2SOO decomposition, leading to O_3^- formation and releasing SO_2. The chemistry of $SO_2 + O_3^-$ has been explored elsewhere. The second mechanism is characterized by SO_2 oxidation and proceeds through formation of a pre-reactive complex that subsequently reacts to form the products by overcoming a relatively low energy barrier. The overall reaction, $O_2^- + SO_2 + O_3 \rightarrow SO_3^- + 2O_2$, is faster and more energetically favorable than the $SO_2 + O_3^- \rightarrow SO_3^- + O_2$ reaction, thereby highlighting the positive role of O_2^- in SO_2 ionic oxidation. Hence, the two reactions may compete in chamber experiments

15 and in the atmosphere.

While for the electron transfer and O_2SOO decomposition process the reaction is hindered by the presence of water, the oxidation reaction is catalysed instead as the rate constant is increased by 6 orders of magnitude with the presence of water. Weighing the rate constants of unhydrated and monohydrated reactions to the equilibrium concentrations of hydrates of corresponding pre-reactive complexes leads to the net rate constant of 4.0×10^{-11} cm^3 molecule^{-1} s^{-1} at 298 K for the oxidation

20 reaction. Hence, this reaction proceeds nearly at collision rate. The main species (SO_3^-) in the end products of the studied reaction has been proved to form both in the atmosphere and in experiments, where it definitely plays a role in atmospheric sulfur chemistry and particle formation. The contribution of this mechanism to the total atmospheric sulfuric acid formation is estimated. The studied reaction further deepens the understanding of ion-induced SO_2 oxidation, with implications in aerosol formation.

25 **Author contributions**

NTT and LD designed the work. NTT performed all calculations and wrote the manuscript. LD edited the manuscript.

Competing interests

The authors declare that they have no conflict of interest.

Acknowledgements

This work was supported by the National Natural Science Foundation of China (21707080, 91644214), the Postdoctoral Science Foundation of China (2017M612276), Shandong Natural Science Fund for Distinguished Young Scholars (JQ201705) and the International Postdoctoral Exchange Fellowship Program. We acknowledge the High-Performance Computing Center
5   of Shandong University for providing the computational resources.

References

Atkinson, R., Baulch, D. L., Cox, R. A., Crowley, J. N., Hampson, R. F., Hynes, R. G., Jenkin, M. E., Rossi, M. J., and Troe, J.: Evaluated kinetic and photochemical data for atmospheric chemistry: Volume I - gas phase reactions of Ox, HOx, NOx and
10   SOx species, Atmos. Chem. Phys., 4, 1461-1738, doi:10.5194/acp-4-1461-2004, 2004.

Bader, R. F. W.: A Bond Path: A Universal Indicator of Bonded Interactions, J. Phys. Chem. A, 102, 7314-7323, doi:10.1021/jp981794v, 1998.

Bork, N., Kurtén, T., Enghoff, M., Pedersen, J. O. P., Mikkelsen, K. V., and Svensmark, H.: Ab initio studies of $O_2^-(H_2O)_n$ and $O_3^-(H_2O)_n$ anionic molecular clusters, $n \leq 12$, Atmos. Chem. Phys., 11, 7133-7142, doi:10.5194/acp-11-7133-2011, 2011.

15   Bork, N., Kurtén, T., Enghoff, M., Pedersen, J. O. P., Mikkelsen, K. V., and Svensmark, H.: Structures and reaction rates of the gaseous oxidation of SO_2 by an $O_3^-(H_2O)_{0-5}$ cluster – a density functional theory investigation, Atmos. Chem. Phys., 12, 3639-3652, doi:10.5194/acp-12-3639-2012, 2012.

Bork, N., Kurtén, T., and Vehkamäki, H.: Exploring the atmospheric chemistry of $O_2SO_3^-$ and assessing the maximum turnover number of ion-catalysed H_2SO_4 formation, Atmos. Chem. Phys., 13, 3695-3703, doi:10.5194/acp-13-3695-2013, 2013.

20   Buttar, D., and Hirst, D. M.: Ab initio quantum chemistry study of the gas-phase reaction of ClO with HO_2, J. Chem. Soc. Faraday Trans., 90, 1811-1817, doi:10.1039/FT9949001811, 1994.

Drougas, E., and Kosmas, A. M.: Computational studies of (HIO_3) isomers and the HO_2+IO reaction pathways, J. Phys. Chem. A, 109, 3887-3892, doi:10.1021/jp044197j, 2005.

Dunning Jr, T. H., Peterson, K. A., and Wilson, A. K.: Gaussian basis sets for use in correlated molecular calculations. X. The
25   atoms aluminum through argon revisited, J. Chem. Phys., 114, 9244-9253, doi:10.1063/1.1367373, 2001.

Ehn, M., Junninen, H., Petäjä, T., Kurtén, T., Kerminen, V.-M., Schobesberger, S., Manninen, H., Ortega, I., Vehkamäki, H., and Kulmala, M.: Composition and temporal behavior of ambient ions in the boreal forest, Atmos. Chem. Phys., 10, 8513-8530, doi:10.5194/acp-10-8513-2010, 2010.

Elm, J., Bilde, M., and Mikkelsen, K. V.: Assessment of Density Functional Theory in Predicting Structures and Free Energies
30   of Reaction of Atmospheric Prenucleation Clusters, J. Chem. Theory Comput., 8, 2071-2077, doi:10.1021/ct300192p, 2012.

Elm, J., Bilde, M., and Mikkelsen, K. V.: Influence of Nucleation Precursors on the Reaction Kinetics of Methanol with the OH Radical, J. Phys. Chem. A, 117, 6695-6701, doi:10.1021/jp4051269, 2013a.

Elm, J., Bilde, M., and Mikkelsen, K. V.: Assessment of binding energies of atmospherically relevant clusters, Phys. Chem. Chem. Phys., 15, 16442-16445, doi:10.1039/c3cp52616j, 2013b.

Enghoff, M., and Svensmark, H.: The role of atmospheric ions in aerosol nucleation–a review, Atmos. Chem. Phys., 8, 4911-4923, doi:10.5194/acp-8-4911-2008, 2008.

5 Enghoff, M. B., Bork, N., Hattori, S., Meusinger, C., Nakagawa, M., Pedersen, J. O. P., Danielache, S., Ueno, Y., Johnson, M. S., Yoshida, N., and Svensmark, H.: An isotopic analysis of ionising radiation as a source of sulphuric acid, Atmos. Chem. Phys., 12, 5319-5327, doi:10.5194/acp-12-5319-2012, 2012.

Fahey, D., Böhringer, H., Fehsenfeld, F., and Ferguson, E.: Reaction rate constants for $O_2^-(H_2O)_n$ ions n= 0 to 4, with O_3, NO, SO_2, and CO_2, J. Chem. Phys., 76, 1799-1805, doi:10.1063/1.443220, 1982.

10 Fehsenfeld, F., and Ferguson, E.: Laboratory studies of negative ion reactions with atmospheric trace constituents, J. Chem. Phys., 61, 3181-3193, doi:10.1063/1.1682474, 1974.

Franchin, A., Ehrhart, S., Leppa, J., Nieminen, T., Gagne, S., Schobesberger, S., Wimmer, D., Duplissy, J., Riccobono, F., Dunne, E. M., Rondo, L., Downard, A., Bianchi, F., Kupc, A., Tsagkogeorgas, G., Lehtipalo, K., Manninen, H. E., Almeida, J., Amorim, A., Wagner, P. E., Hansel, A., Kirkby, J., Kuerten, A., Donahue, N. M., Makhmutov, V., Mathot, S., Metzger, A.,

15 Petaja, T., Schnitzhofer, R., Sipila, M., Stozhkov, Y., Tome, A., Kerminen, V. M., Carslaw, K., Curtius, J., Baltensperger, U., and Kulmala, M.: Experimental investigation of ion-ion recombination under atmospheric conditions, Atmos. Chem. Phys., 15, 7203-7216, doi:10.5194/acp-15-7203-2015, 2015.

Fukui, K.: The path of chemical reactions-the IRC approach, Acc. Chem. Res., 14, 363-368, doi:10.1021/ar00072a001, 1981.

Georgievskii, Y., and Klippenstein, S. J.: Long-range transition state theory, J. Chem. Phys., 122, 10.1063/1.1899603, 2005.

20 Greenwald, E. E., North, S. W., Georgievskii, Y., and Klippenstein, S. J.: A two transition state model for radical-molecule reactions: A case study of the addition of OH to C_2H_4, J. Phys. Chem. A, 109, 6031-6044, doi:10.1021/jp058041a, 2005.

Harris, E., Sinha, B., van Pinxteren, D., Tilgner, A., Fomba, K. W., Schneider, J., Roth, A., Gnauk, T., Fahlbusch, B., Mertes, S., Lee, T., Collett, J., Foley, S., Borrmann, S., Hoppe, P., and Herrmann, H.: Enhanced Role of Transition Metal Ion Catalysis During In-Cloud Oxidation of SO_2, Science, 340, 727-730, doi:10.1126/science.1230911, 2013.

25 Henkelman, G., Arnaldsson, A., and Jónsson, H.: A fast and robust algorithm for Bader decomposition of charge density, Comput. Mater. Sci., 36, 354-360, https://doi.org/10.1016/j.commatsci.2005.04.010, 2006.

Hvelplund, P., Pedersen, J. O. P., Stochkel, K., Enghoff, M. B., and Kurten, T.: Experimental studies of the formation of cluster ions formed by corona discharge in an atmosphere containing SO_2, NH_3, and H_2O, Int. J. Mass Spectrom., 341, 1-6, doi:10.1016/j.ijms.2013.03.001, 2013.

30 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagne, S., Ickes, L., Kuerten, A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G., Wimmer, D., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud, W., Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R., Makhmutov, V., Mathot, S., Mikkila, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A., Pereira, P., Petaja, T., Schnitzhofer, R., Seinfeld, J. H., Sipila,

M., Stozhkov, Y., Stratmann, F., Tome, A., Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U., and Kulmala, M.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation, Nature, 476, 429-433, doi:10.1038/nature10343, 2011.

Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M., Yan, C., Almeida, J., Tröstl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A., Bernhammer, A.-K., Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X., Craven, J., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim, J., Krapf, M., Kürten, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Peräkylä, O., Piel, F., Petäjä, T., Praplan, A. P., Pringle, K., Rap, A., Richards, N. A. D., Riipinen, I., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Scott, C. E., Seinfeld, J. H., Sipilä, M., Steiner, G., Stozhkov, Y., Stratmann, F., Tomé, A., Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R., Baltensperger, U., Kulmala, M., Carslaw, K. S., and Curtius, J.: Ion-induced nucleation of pure biogenic particles, Nature, 533, 521-526, doi:10.1038/nature17953, 2016.

Klippenstein, S. J., Khundkar, L. R., Zewail, A. H., and Marcus, R. A.: Application of unimolecular reaction-rate theory for highly flexible transition-states to the dissociation of NCNO into NC and NO, J. Chem. Phys., 89, 4761-4770, doi:10.1063/1.455670, 1988.

Kozuch, S., and Martin, J. M. L.: "Turning Over" Definitions in Catalytic Cycles, ACS Catal., 2, 2787-2794, doi:10.1021/cs3005264, 2012.

Kupiainen-Määttä, O., Olenius, T., Kurtén, T., and Vehkamäki, H.: CIMS sulfuric acid detection efficiency enhanced by amines due to higher dipole moments: a computational study, J. Phys. Chem. A, 117, 14109-14119, doi:10.1021/jp4049764, 2013.

Larson, L. J., Kuno, M., and Tao, F.-M.: Hydrolysis of sulfur trioxide to form sulfuric acid in small water clusters, The Journal of Chemical Physics, 112, 8830-8838, 2000.

Mauldin III, R. L., Berndt, T., Sipilä, M., Paasonen, P., Petäjä, T., Kim, S., Kurtén, T., Stratmann, F., Kerminen, V.-M., and Kulmala, M.: A new atmospherically relevant oxidant of sulphur dioxide, Nature, 488, 193-196, doi:10.1038/nature11278, 2012.

Möhler, O., Reiner, T., and Arnold, F.: The formation of SO_5^- by gas phase ion–molecule reactions, J. Chem. Phys., 97, 8233-8239, doi:10.1063/1.463394, 1992.

Nagato, K., Kim, C. S., Adachi, M., and Okuyama, K.: An experimental study of ion-induced nucleation using a drift tube ion, mobility spectrometer/mass spectrometer and a cluster-differential mobility analyzer/Faraday cup electrometer, J. Aerosol Sci., 36, 1036-1049, doi:10.1016/j.jaerosci.2004.12.006, 2005.

Nieminen, T., Manninen, H., Sihto, S.-L., Yli-Juuti, T., Mauldin III, R. L., Petaja, T., Riipinen, I., Kerminen, V.-M., and Kulmala, M.: Connection of sulfuric acid to atmospheric nucleation in boreal forest, Environ. Sci. Technol., 43, 4715-4721, doi: 10.1021/es803152j, 2009.

Noziere, B., Ekstrom, S., Alsberg, T., and Holmstrom, S.: Radical-initiated formation of organosulfates and surfactants in atmospheric aerosols, Geophys. Res. Lett., 37, doi:10.1029/2009gl041683, 2010.

Peng, C., Ayala, P. Y., Schlegel, H. B., and Frisch, M. J.: Using redundant internal coordinates to optimize equilibrium geometries and transition states, J. Comput. Chem., 17, 49-56,, 1996.

Purvis, G. D., and Bartlett, R. J.: A full coupled-cluster singles and doubles model - the inclusion of disconnected triples, J. Chem. Phys., 76, 1910-1918, doi:10.1063/1.443164, 1982.

5 Rothe, E. W., Tang, S. Y., and Reck, G. P.: Measurement of electron affinities of O_3, SO_2, and SO_3 by collisional ionization, J. Chem. Phys., 62, 3829-3831, doi:10.1063/1.430941, 1975.

Schindelka, J., Iinuma, Y., Hoffmann, D., and Herrmann, H.: Sulfate radical-initiated formation of isoprene-derived organosulfates in atmospheric aerosols, Faraday Discuss., 165, 237-259, doi:10.1039/c3fd00042g, 2013.

Schweitzer, C., and Schmidt, R.: Physical mechanisms of generation and deactivation of singlet oxygen, Chem. Rev., 103, 1685-
10 1757, doi:10.1021/cr010371d, 2003.

Seinfeld, J. H., and Pandis, S. N.: Atmospheric Chemistry and Physics: From Air Pollution to Climate Change, 3rd Edition, John Wiley & Sons, Inc., 2016.

Sipila, M., Berndt, T., Petaja, T., Brus, D., Vanhanen, J., Stratmann, F., Patokoski, J., Mauldin III, R. L., Hyvarinen, A.-P., Lihavainen, H., and Kulmala, M.: The Role of Sulfuric Acid in Atmospheric Nucleation, Science, 327, 1243-1246,
15 doi:10.1126/science.1180315, 2010.

Su, T., and Chesnavich, W. J.: Parametrization of the ion-polar molecule collision rate-constant by trajectory calculations, J. Chem. Phys., 76, 5183-5185, doi:10.1063/1.442828, 1982.

Surratt, J. D., Lewandowski, M., Offenberg, J. H., Jaoui, M., Kleindienst, T. E., Edney, E. O., and Seinfeld, J. H.: Effect of acidity on secondary organic aerosol formation from isoprene, Environ. Sci. Technol., 41, 5363-5369, doi:10.1021/es0704176, 2007.

20 Surratt, J. D., Gomez-Gonzalez, Y., Chan, A. W. H., Vermeylen, R., Shahgholi, M., Kleindienst, T. E., Edney, E. O., Offenberg, J. H., Lewandowski, M., Jaoui, M., Maenhaut, W., Claeys, M., Flagan, R. C., and Seinfeld, J. H.: Organosulfate formation in biogenic secondary organic aerosol, J. Phys. Chem. A, 112, 8345-8378, doi:10.1021/jp802310p, 2008.

Svensmark, H., Pedersen, J. O. P., Marsh, N. D., Enghoff, M. B., and Uggerhoj, U. I.: Experimental evidence for the role of ions in particle nucleation under atmospheric conditions, Proc. R. Soc. A, 463, 385-396, doi:10.1098/rspa.2006.1773, 2007.

25 Tang, W., Sanville, E., and Henkelman, G.: A grid-based Bader analysis algorithm without lattice bias, J. Phys. Condens. Matter, 21, doi:10.1088/0953-8984/21/8/084204, 2009.

Tsona, N., Bork, N., and Vehkamäki, H.: Exploring the chemical fate of the sulfate radical anion by reaction with sulfur dioxide in the gas phase, Atmos. Chem. Phys., 15, 495-503, doi:10.5194/acp-15-495-2015, 2015.

Tsona, N. T., Bork, N., and Vehkamäki, H.: On the gas-phase reaction between SO_2 and $O_2^-(H_2O)_{0-3}$ clusters – an ab initio study, Phys.
30 Chem. Chem. Phys., 16, 5987-5992, doi:10.1039/c3cp54715a, 2014.

Tsona, N. T., Bork, N., Loukonen, V., and Vehkamäki, H.: A Closure Study of the Reaction between Sulfur Dioxide and the Sulfate Radical Ion from First-Principles Molecular Dynamics Simulations, J. Phys. Chem. A, 120, 1046-1050, doi: 10.1021/acs.jpca.5b12395, 2016.

Tsona, N. T., Li, J., and Du, L.: From O_2^--Initiated SO_2 oxidation to sulfate formation in the gas-phase, J. Phys. Chem. A, 122, 5781-5788, doi:10.1021/acs.jpca.8b03381, 2018.

Vereecken, L., Harder, H., and Novelli, A.: The reaction of Criegee intermediates with NO, RO_2, and SO_2, and their fate in the atmosphere, Phys. Chem. Chem. Phys., 14, 14682-14695, doi:10.1039/c2cp42300f, 2012.

5   Wagner, R., Yan, C., Lehtipalo, K., Duplissy, J., Nieminen, T., Kangasluoma, J., Ahonen, L. R., Dada, L., Kontkanen, J., Manninen, H. E., Dias, A., Amorim, A., Bauer, P. S., Bergen, A., Bernhammer, A. K., Bianchi, F., Brilke, S., Mazon, S. B., Chen, X., Draper, D. C., Fischer, L., Frege, C., Fuchs, C., Garmash, O., Gordon, H., Hakala, J., Heikkinen, L., Heinritzi, M., Hofbauer, V., Hoyle, C. R., Kirkby, J., Kürten, A., Kvashnin, A. N., Laurila, T., Lawler, M. J., Mai, H., Makhmutov, V., Mauldin III, R. L., Molteni, U., Nichman, L., Nie, W., Ojdanic, A., Onnela, A., Piel, F., Quéléver, L. L. J., Rissanen, M. P.,

10   Sarnela, N., Schallhart, S., Sengupta, K., Simon, M., Stolzenburg, D., Stozhkov, Y., Tröstl, J., Viisanen, Y., Vogel, A. L., Wagner, A. C., Xiao, M., Ye, P., Baltensperger, U., Curtius, J., Donahue, N. M., Flagan, R. C., Gallagher, M., Hansel, A., Smith, J. N., Tomé, A., Winkler, P. M., Worsnop, D., Ehn, M., Sipilä, M., Kerminen, V. M., Petäjä, T., and Kulmala, M.: The role of ions in new particle formation in the CLOUD chamber, Atmos. Chem. Phys., 17, 15181-15197, doi:10.5194/acp-17-15181-2017, 2017.

15   Welz, O., Savee, J. D., Osborn, D. L., Vasu, S. S., Percival, C. J., Shallcross, D. E., and Taatjes, C. A.: Direct kinetic measurements of Criegee intermediate (CH_2OO) formed by reaction of CH_2I with O_2, Science, 335, 204-207, doi:10.1126/science.1213229, 2012.

Yan, C., Dada, L., Rose, C., Jokinen, T., Nie, W., Schobesberger, S., Junninen, H., Lehtipalo, K., Sarnela, N., Makkonen, U., Garmash, O., Wang, Y., Zha, Q., Paasonen, P., Bianchi, F., Sipilä, M., Ehn, M., Petäjä, T., Kerminen, V. M., Worsnop, D. R.,

20   and Kulmala, M.: The role of H2SO4-NH3 anion clusters in ion-induced aerosol nucleation mechanisms in the boreal forest, Atmos. Chem. Phys., 18, 13231-13243, doi:10.5194/acp-18-13231-2018, 2018.

Yu, F.: From molecular clusters to nanoparticles: second-generation ion-mediated nucleation model, Atmos. Chem. Phys., 6, 5193-5211, doi:10.5194/acp-6-5193-2006, 2006.

Yu, F. Q., and Turco, R. P.: Ultrafine aerosol formation via ion-mediated nucleation, Geophys. Res. Lett., 27, 883-886,

25   doi:10.1029/1999gl011151, 2000.

Yu, F. Q., and Turco, R. P.: From molecular clusters to nanoparticles: Role of ambient ionization in tropospheric aerosol formation, J. Geophys. Res.-Atmos., 106, 4797-4814, doi:10.1029/2000jd900539, 2001.

Zhao, Y., and Truhlar, D. G.: The M06 suite of density functionals for main group thermochemistry, thermochemical kinetics, noncovalent interactions, excited states, and transition elements: two new functionals and systematic testing of four M06-class

30   functionals and 12 other functionals, Theor. Chem. Account., 120, 215-241, doi:10.1007/s00214-007-0310-x, 2008.

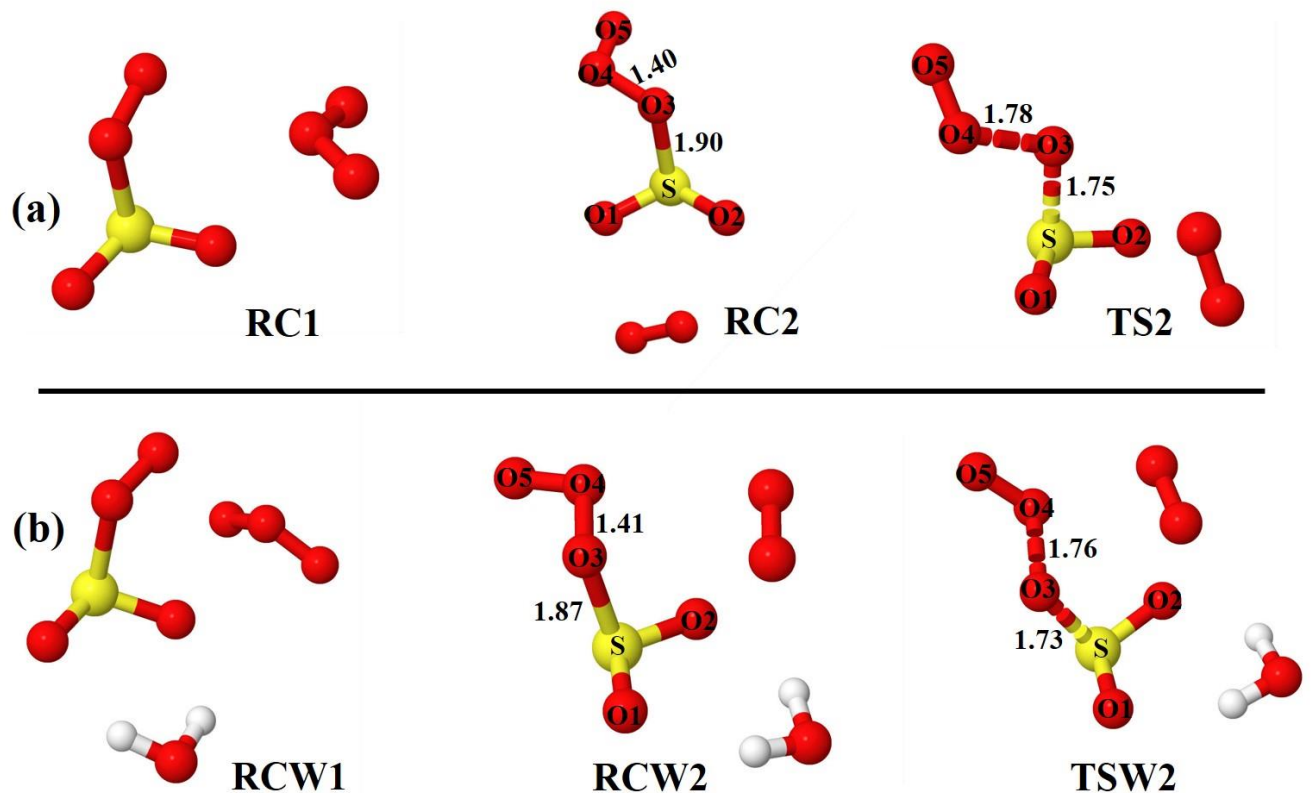



Figure 1: Optimized structures of the most stable intermediates in the O_2SOO^- + O_3 reaction in the absence (a) and in the presence (b) of a single water molecule. Optimizations were performed at the M06-2X/aug-cc-pVTZ level of theory. Lengths (in Å) of some descriptive bonds are indicated. The color coding is yellow for sulfur, red for oxygen and white for hydrogen.

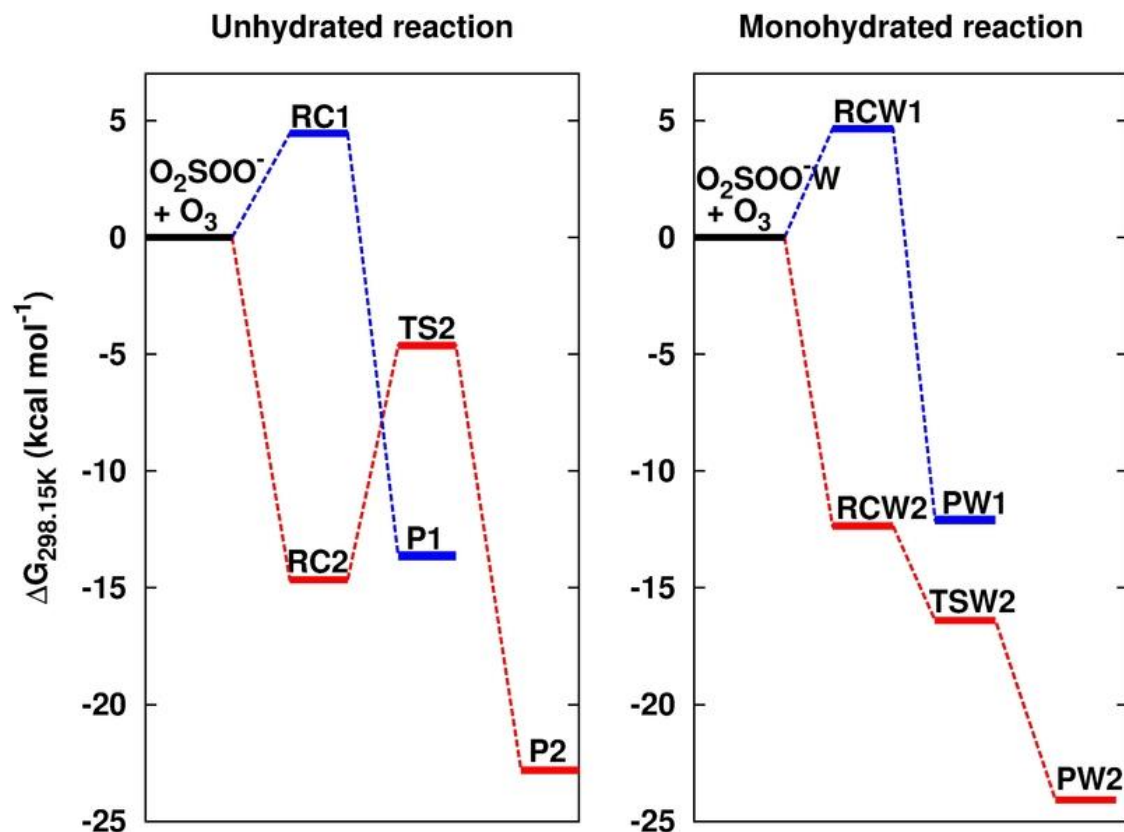



Figure 2: Formation Gibbs free energies of the most stable intermediate species in the $O_2SOO^- + O_3$ reaction in the absence and in the presence of water. "W" is the shorthand notation for water. RC1, RC2, TS2, RCW1, RCW2, and TSW2 structures are shown in Fig. 1. P1 = $O_2 + SO_2 + O_3^-$, P2 = $SO_3^- + 2O_2$, PW1 = $O_2 + SO_2 + O_3^- + H_2O$ and PW2 = $SO_3\cdots H_2O + 2O_2$. Calculations were performed at the CCSD(T)/aug-cc-pVTZ//M06-2X/aug-cc-pVTZ level of theory.

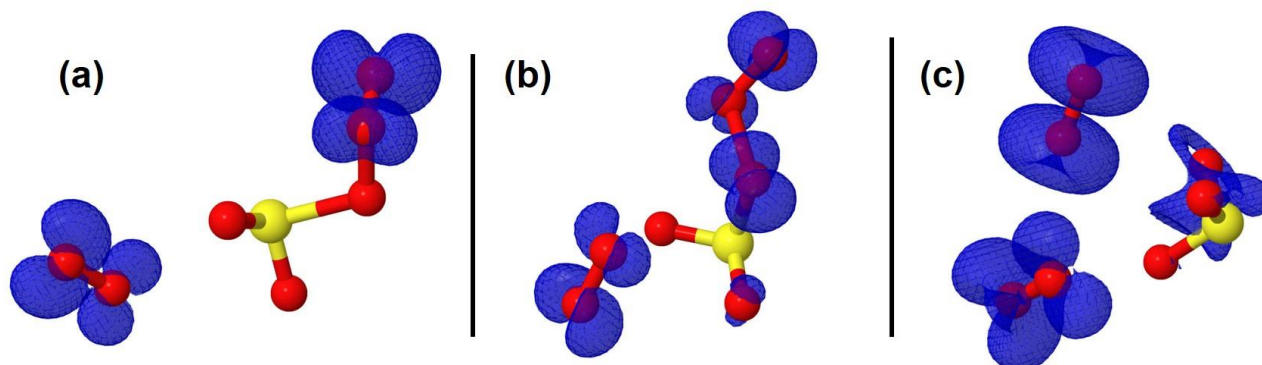



Figure 3: Representation of the spin density (in blue color) on intermediate structures in the $O_2SOO^- + O_3$ reaction. The spin density clearly indicates that the extra electron is progressively distributed over all the atoms from (a) the pre-reactive complex through (b) the transition state to (c) the product complex.

Table 1: Electronic energies (ΔE), enthalpies (ΔH_{298K}) and Gibbs free energies (ΔG_{298K}) of the different states in the $O_2SOO^- + O_3$ reaction both in the absence and in the presence of water, calculated relative to the energy of initial reactants at the CCSD(T)/aug-cc-pVTZ//M06-2X/aug-cc-pVTZ level of theory.

Species	ΔE	ΔH_{298K}	ΔG_{298K}
Unhydrated reaction			
$O_2SOO^- + O_3$	0	0	0
$RC1=O_3\cdots O_2SOO^-$	-5.1	-3.9	4.5
$RC2=O_2\cdots O_2S–O_3^-$	-21.9	-21.0	-14.7
TS2	-11.6	-11.9	-4.6
$SO_3^- + 2O_2$	-13.3	-13.9	-22.8
$O_2 + SO_2 + O_3^-$	-1.7	-3.2	-13.6
Monohydrated reaction			
$O_2SOO^-\cdots H_2O + O_3$	0	0	0
$RCW1=O_3\cdots O_2SOO^-\cdots H_2O$	-4.6	-3.4	4.7
$RCW2=O_2\cdots O_2S–O_3^-\cdots H_2O$	-21.9	-20.9	-12.4
TSW2	-25.3	-25.4	-16.4
$SO_3^-\cdots H_2O + 2O_2 + H_2O$	-14.2	-14.9	-24.1
$O_2 + SO_2 + O_3^- + H_2O$	10.3	7.1	-12.1