

1 **Source contributions to sulfur and nitrogen deposition – an HTAP II multi-**
2 **model study on hemispheric transport**

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21 **Abstract.** With the rising anthropogenic emissions from human activities, elevated concentrations
22 of air pollutants have been detected in the hemispheric air flows in recent years, aggravating the
23 regional air pollution and deposition issues. However, regional contributions of hemispheric air
24 pollution to deposition at global scale have been given little attention in the literature. In this light,
25 we assess the impact of hemispheric transport on sulfur (S) and nitrogen (N) deposition for 6 world
26 regions: North America (NA), Europe (EU), South Asia (SA), East Asia (EA), Middle East (ME)
27 and Russia (RU) in 2010, by using the multi-model ensemble results from the 2nd phase of Task
28 Force Hemispheric Transport of Air Pollution (HTAP II) with 20% emission perturbation
29 experiments. About 27-58%, 26-46% and 12-23% of local S, NO_x and NH₃ emissions are
30 transported and removed by deposition outside of the source regions annually, with seasonal
31 variation of 5% more in winter and 5% less in summer. The 20% emission reduction in source
32 regions could affect 1-10% of deposition on foreign continental regions and 1-14% on foreign
33 coastal regions and the open ocean. Significant influences are found from NA to the North Atlantic
34 Ocean (2-14%), and from EA to the North Pacific Ocean (4-10%) and western NA (4-6%) with
35 20% emission reduction in source regions. The impact on deposition caused by short-distance

36 transport between neighbouring regions (i.e. EU and RU) occurs throughout the whole year
37 (slightly stronger in winter), while the long-range transport (i.e. from EA to NA) mainly takes
38 place in spring and fall, which is consistent with the seasonality found for hemispheric transport
39 of air pollutants. Deposition in emission intensive regions such as EA is dominated (~80%) by
40 own region emission, while deposition in low emission intensity regions such as RU is almost
41 equally affected by foreign emission (40-60%) and own region emission. We also find that
42 deposition on the coastal regions or the near-coastal open ocean is twice more sensitive to
43 hemispheric transport than non-coastal continental regions, especially for regions (i.e. west coast
44 of NA) in the downwind location of emission source regions. This study highlights the significant
45 impact of hemispheric transport on the deposition of coastal regions, the open ocean and low
46 emission intensity regions. Further research is proposed for improving ecosystem and human
47 health, with regards to the enhanced hemispheric transport.

48 **1 Introduction**

49 The increasing consumption of energy by human activities has largely increased the deposition of
50 nitrogen (N) over the terrestrial and marine ecosystem (Kim et al., 2011; Galloway et al., 2008;
51 Duce et al., 2008). The impact is estimated to continue increasing in the near future (Bleeker et al.,
52 2011; Lamarque et al., 2013; Kanakidou et al., 2016; Paulot et al., 2013; Lamarque et al., 2005;
53 Bian et al., 2017). The NO_x emission has increased by about 10 Tg(N) from 2001 to 2010, due to
54 large increase in Asia regions (Tan et al., 2018), but the recent studies report a turning point for
55 Chinese NO_x emission in 2011 (Li et al., 2017; Liu et al., 2016). On the other hand, the global
56 sulfur (S) emission has declined by about 5 Tg(S) from 2000 to 2010 (Tan et al., 2018). The global
57 fossil fuel SO₂ emission has a decreasing trend since 1980 owing to the significant decline in SO₂
58 emission from Europe (EU) and U.S. (Chin et al., 2014). The SO₂ emission in China experiences
59 increases from 2000-2005 due to energy consumption and decreases after 2006 thanks to the
60 implementation of Flue-Gas Desulphurization system. Deposition supplies the ecosystem with
61 nutrients, but too much deposition could cause various adverse impacts on the environment,
62 including acidification and eutrophication of the forest and waterbody (Bouwman et al., 2002;
63 Bergstrom and Jansson, 2006; Dentener et al., 2006; Phoenix et al., 2006), soil acidification that
64 slows down the crop production (Guo et al., 2010; Janssens et al., 2010) and even destroying the

65 plant biodiversity (Bobbink et al., 2010; Clark and Tilman, 2008). The prevention and control of
66 exceeding deposition have become a growing worldwide concern.

67 Hemispheric transport of air pollutants is found to aggravate the regional air pollution
68 issues (Wild and Akimoto, 2001; Sudo and Akimoto, 2007; Fu et al., 2012; Fiore et al., 2009) as
69 well as enlarge the local deposition burden (Glotfelty et al., 2014; Sanderson et al., 2008). The
70 mass of aerosols arriving at the North American coasts is comparable to that emitted domestically
71 (Yu et al., 2012). Air pollution from Asia contributes to the PM_{2.5} concentration in western U.S.
72 by 1.5 µg m⁻³ (Tao et al., 2016), the O₃ concentration by 3-10 ppbv (Zhang et al., 2009; Zhang et
73 al., 2008; Yienger et al., 2000; Reidmiller et al., 2009; Jacob et al., 1999; Brown-Steiner and Hess,
74 2011) and the peroxyacyl nitrate (PAN) concentration by 26 ppbv (Berntsen et al., 1999) in spring.
75 The long-range transport of air pollution from North America (NA) is estimated to contribute by
76 3-5 ppb (7-11%) to the O₃ concentration in EU annually (Auvray and Bey, 2005; Guerova et al.,
77 2006; Derwent et al., 2004; Li et al., 2002) and the increment can reach 25-28 ppbv during
78 particular events (Guerova et al., 2006). European outflow affects the surface O₃ concentration in
79 western China by 2-6 ppbv in spring and summer (Li et al., 2014) and North Africa by 5-20 ppbv
80 in summer (Duncan et al., 2008; Duncan and Bey, 2004). The study by Yu et al. (2013) found that
81 the long-range transport contributes by 6-16% and 22-40% to aerosol optical depth and direct
82 radiative forcing in 4 regions including NA, EU, East Asia (EA) and South Asia (SA). Recent
83 studies have reported an increasing trend in the hemispheric transport of air pollution from Asia to
84 NA from mid-1980s to late-2000s. The Asian plume has contributed ~10 ppbv (30%) to the O₃
85 concentration over western NA from mid-1980s to mid-2000s (Jaffe et al., 2003; Parrish et al.,
86 2004), with an annual increase of 0.34-0.50 ppbv O₃ (Parrish et al., 2009). More recent study showed
87 the contribution is about 5-7 ppbv O₃ in 2006 with an annual increase rate of 1-2 ppb O₃ since 2000
88 (Zhang et al., 2008). The trend well agreed with the rapid growth of Asian emission (Richter et al.,
89 2005; Lu et al., 2010; Verstraeten et al., 2015; Zhang et al., 2007; van der A et al., 2006; van der
90 A et al., 2008).

91 Compared to the impact on air pollution, the impact of hemispheric transport on deposition
92 hasn't been fully studied. Arndt and Carmichael (1995) developed a source-receptor (S-R)
93 relationship for S deposition among the Asia regions in early 1990s. Zhang et al. (2012) found that
94 foreign anthropogenic emission contributes to 6% and 8% of the oxidized nitrogen (NO_y) and
95 reduced nitrogen (NH_x) deposition in contiguous U.S., respectively. A systematic study by

96 Sanderson et al. (2008) shed light on the impact of long-range transport on NO_y deposition at
97 global scale. The study used the model ensemble results from the 1st phase of Task Force
98 Hemispheric Transport of Air Pollution (HTAP I) to calculate the S-R relationship for NO_y
99 deposition in 2001 among 4 regions: EU, NA, SA and EA. Results showed that about 12-24% of
100 the NO_x emission was transported and deposited out of source regions. About 3-10% of the
101 emission was deposited on the other 3 regions and affects their deposition by about 1-3%. However,
102 these studies focused on the emission intensive regions, where the foreign disturbance could be
103 relatively small compared to huge own region emission. The foreign impact on low emission
104 intensity regions was not evaluated in the same detail. Furthermore, both the magnitude and spatial
105 distribution of S and N emission and deposition have been changed considerably during the last
106 10 years (2001-2010) (Tan et al., 2018). It is necessary to update the S-R relationship for more
107 recent years with regards to these changes.

108 To explore these questions, this study assesses the impact of hemispheric transport of S,
109 NO_x and NH_3 emissions on S and N deposition, with multi-model ensemble results from the 2nd
110 phase of HTAP (HTAP II). Additional to the 4 regions: NA, EU, SA, EA used in Sanderson's
111 study for HTAP I, we include 2 more regions: Middle East (ME) and Russia, Belarussia, Ukraine
112 (RU) in this study. These two regions have low S and N emissions relative to their areal extent,
113 but are located close to the high emission regions such as EU, SA and EA. We calculate the amount
114 of deposition brought by hemispheric transport by comparing model results between the base case
115 and the 20% emission perturbation cases. The experimental design is described in Section 2.
116 Section 3 is the result part and has 3 subsections. We explore the following questions:

- 117 1) What fraction (percentage) of the S or N emissions is transported and deposited outside of each
118 source region? What is the fraction of emission that finally deposited on the other 5 receptor
119 regions and oceans? What is the seasonality of the exported fraction?
- 120 2) As receptor regions, how much deposition is brought by hemispheric transport? What is the
121 impact on local deposition? Is there any seasonality for this impact?
- 122 3) For each region, what are the contributions on deposition of hemispheric transport from foreign
123 regions and of own region emission? In line with the analysis for other pollutants, to this
124 purpose we evaluate the so-called response to extra-regional emission reduction (RERER)
125 metric. We also discuss the own region and foreign impact on the coastal regions specifically.
126 The inter-model variations are also illustrated.

127 Section 4 is a summary of the findings in this study and some suggestions for future study.

128 **2 Methodology**

129 **2.1 HTAP II and experiment set-up**

130 The HTAP was created in 2004 under the Convention on Long-range Transboundary Air Pollution
131 (CLRTAP). It involves efforts from international scientists aiming at understanding the
132 hemispheric transport of air pollutants and its impact on regional and global air quality, public
133 health and near-term climate change. Until now, two phases of HTAP experiments have been
134 conducted successfully. The HTAP I involved more than 20 models from international modeling
135 groups, with 2001 as the base year for modeling studies. A comprehensive report of the major
136 findings of HTAP I was released in 2010 and could be downloaded from <http://www.htap.org/>.
137 The HTAP II was launched in 2012, with 2010 as the base simulation year. HTAP II required all
138 models to use the same prescribed anthropogenic emission instead of using the best estimates of
139 emission by each modeling group as HTAP I. This facilitates an inter-model comparison between
140 models. In addition, HTAP II had a refined definition for the boundaries of regions, which enabled
141 an update in the S-R relationships for air pollutants and deposition among regions.

142 This study uses the ensemble of 11 global models from HTAP II (including CAM-Chem,
143 CHASER_re1, CHASER_t106, EMEP_rv48, GEMMACH, GEOS5, GEOSCHEMAJOINT,
144 OsloCTM3v.2, GOCARTv5, SPRINTARS and C-IFS_v2). A detailed description of the
145 experiment set-up could be found in Galmarini et al. (2017). The S deposition includes SO₂
146 deposition and SO₄²⁻ deposition. N deposition is categorized to NO_y and NH_x deposition. NO_y
147 deposition is a sum of all oxidized N except N₂O, including NO₂, HNO₃, NO₃⁻, PAN and other
148 organic nitrates than PAN (Orgn). NH_x deposition includes NH₃ deposition and NH₄⁺ deposition.
149 To form the multi-model ensemble, we re-grid all models to a uniformed horizontal resolution of
150 0.1° × 0.1°. We use the multi-model mean value (MMM) of all models to present the ensemble
151 results, a procedure which has been proven previously to have a better agreement with observations
152 than single model result (Dentener et al., 2006; Tan et al., 2018). The multi-model mean values of
153 the compositions of S or N deposition are calculated separately and then combined to compute the
154 total S or N deposition. More details can be found in Tan et al. (2018).

155 2.2 Simulation scenarios

156 The base simulation uses anthropogenic emissions in 2010 (Janssens-Maenhout et al.,
157 2015), which is called “base case” in this study. The MMM performance on wet deposition has
158 been evaluated with observations from National Atmospheric Deposition Program (NADP)
159 (<http://nadp.sws.uiuc.edu/>, last access: 6 April 2018) for NA, European Monitoring and Evaluation
160 Programme (EMEP) CCC reports (<http://www.nilu.no/projects/ccc/reports.html>, last access: 6
161 April 2018) for EU and Acid Deposition Monitoring Network in East Asia (EANET)
162 (<http://www.eanet.asia/>, last access: 6 April 2018) for EA in the previous study of Tan et al. (2018).
163 Following are some brief results of the evaluation results. Modeled gas phase SO₂ wet deposition
164 and aerosol SO₄²⁻ wet deposition are evaluated with observed SO₄²⁻ wet deposition. 76% of the
165 stations of all networks are predicted within ±50% of observation. Negative model biases (-20%)
166 are found at some East Asian stations. Modeled gas phase HNO₃ wet deposition and aerosol NO₃⁻
167 wet deposition are compared with observed NO₃⁻ wet deposition. 83% of the stations of all
168 networks are within ±50% of observation. The European and Southeast Asian stations with high
169 observed NO₃⁻ wet deposition are somewhat underestimated. Modeled gas phase NH₃ wet
170 deposition and aerosol NH₄⁺ wet deposition are compared with observed NH₄⁺ deposition. 81% of
171 modeled NH₄⁺ wet deposition at stations of all networks are within ±50% of observation. A general
172 underestimation is found in modeled NH₄⁺ wet deposition, especially at East Asian stations. In
173 terms of dry deposition, due to the lack of directly measured data, we compare the modeled dry
174 deposition with inferential data from the Clean Air Status and Trends Network (CASTNET) over
175 U.S. The CASTNET data is calculated with observed aerosol concentration and modeled dry
176 deposition velocity, therefore it might have high uncertainty in data quality. Comparison shows
177 that the modeled dry deposition is generally higher than the CASTNET inferential data by a factor
178 of 1-2. This is a common feature of many global and regional models (WMO, 2017). According
179 to the analysis, the model bias for dry deposition mainly comes from the model over-prediction of
180 air pollutant concentration. The CASTNET sites are generally located in remote regions with
181 relatively lower air pollutant concentrations than urban regions, but the models fail to represent
182 this characteristic with coarse spatial resolution (Tan et al., 2018).

183 HTAP II defines the boundaries of 17 regions as shown in Fig. 1. The emission perturbation
184 experiments are conducted separately for 6 regions (regions with color in Fig.1) with high priority:
185 NA, EU, SA, EA, ME and RU. In the perturbation experiments, the anthropogenic emissions

186 (including NO_x, SO₂, NH₃, VOC, CO and PM) of a specific region are reduced by 20% from the
187 amounts in the base case simulation, while the emissions in other regions keep constant. In addition,
188 a global perturbation experiment referred as “GLO” is conducted with 20% reduction of global
189 anthropogenic emissions. We estimate the impact of hemispheric transport on deposition by
190 comparing the model results under perturbation experiments with those under base case simulation.
191 In order to validate the quality of model outputs, we check the mass balance between emission and
192 deposition at global scale. The mass balance check for base case simulation is shown in Tan et al.
193 (2018), therefore we show the mass balance for perturbation experiments in this study. We
194 compare the global total amounts of changes of deposition (Δ Depo) with changes of emissions (Δ
195 Emis) for all perturbation cases (Table S1). Models are excluded if their global Δ Depo values fall
196 outside the range of $\pm 20\%$ of their global Δ Emis. According to our results, the amounts of Δ Depo
197 are almost equivalent to Δ Emis for all perturbation cases except the NH_x deposition under EA
198 case. The Δ Depo of NH_x deposition under EA case is not available due to lack of model results
199 for the NH₄⁺ wet deposition under 20% emission perturbation in EA.

200 **3 Results**

201 **3.1 Export of S and N emissions from source regions**

202 This section studies the export of S and N emissions and oxidation products from source regions.
203 Table 1 shows the S-R relationship of S and N deposition among the 6 regions. The numbers are
204 the sensitivity ($SEN_{r \rightarrow s}$) of deposition in the receptor/source regions to emission changes in the
205 source regions (Sanderson et al., 2008). The metric is calculated as Δ Depo in the receptor/source
206 regions divided by Δ Emis in the source regions following Eq. (1).

$$207 \quad SEN_{r \rightarrow s} = \frac{\Delta Depo (r/s)}{\Delta Emis (s)} \times 100\% \quad (1)$$

208 where s is the source region and r is the receptor region. Δ Depo (r/s) is the deposition change in
209 the receptor/source regions, Δ Emis (S) is the emission change in the source regions. This value
210 indicates the fraction of emission of source regions that is deposited locally or exported to foreign
211 regions.

212 The numbers outside of the parenthesis in Table 1 are for coastal and non-coastal regions
213 together and the numbers in the parenthesis are specifically for coastal regions (defined in Fig. 1).

214 “Others” means the other regions in the world than the 6 regions (white color in Fig.1). The NA
215 region has 69% of its S emission deposited within itself, including 9% deposited on its coastal
216 region. The remaining 31% is exported to the other regions, mostly to the “Others” and less than
217 3% is deposited on the other 5 regions (EU, SA, EA, ME and RU). A relatively large fraction (14
218 %) of European S emissions are exported to RU region. Other major pathways of export of S
219 emissions/reaction products are from SA to EA (9%), from EA to RU (5%) and from RU to EU
220 (7%) and EA (5%). ME has considerable percentages of S emission exported to its nearby regions
221 such as SA (8%), EA (5%) and RU (5%). The S-R relationship of NO_y deposition is similar to that
222 of S deposition, except that EU and ME have 66% and 54% of NO_x emissions deposited within
223 the source region, which are 6% and 12% higher than those of S emissions, likely due to somewhat
224 longer lifetimes of S emission compared to NO_x emission and the large average emission altitude
225 of S emissions. In terms of NH_x deposition, about 20% more NH_3 emission is deposited within the
226 source regions (except ME) compared with S and NO_x emission, due to its short lifetime in the
227 atmosphere.

228 The seasonal variations of the export of S and N emissions by source regions is shown in
229 Fig. S1. In terms of S emission, there is 5-10% of seasonal differences around the annual average
230 export fractions for all regions except SA. SA exports almost half of its S emission in spring, which
231 is twice the numbers in summer (20%) and fall (25%), related to the specific dry period and
232 monsoon circulation. The seasonal export fractions of NO_x and NH_3 emission are similar to that
233 of S emission in the pattern, but generally lower in values in all seasons. Generally, the source
234 regions export the highest percentage of their emissions in winter and spring and lowest in summer.
235 More proficient oxidation chemistry in summer results in more soluble component, and local
236 weather systems, especially the episodic of precipitation has large influence on this seasonality.
237 For most continental regions, the wet deposition accounts for 50-70% of total deposition (Tan et
238 al., 2018; Vet et al., 2014; Dentener et al., 2006). Therefore, local precipitation plays an important
239 role in the local pollution removal process. On the other hand, for regions with low local
240 precipitation like ME, the percentage of emission removed within own region would be lower than
241 the other regions. In addition, the strong westerly winds in winter and spring favor the hemispheric
242 transport for regions in mid-latitudes of the North Hemisphere. While the rapid vertical convection
243 in summer slows down the zonal transport of air flows and accelerates the local removal process.

244 A comparison is conducted with previous studies. Bey et al. (2001) estimated that 70% of
245 emitted NO_x from Asia is lost within its boundary by deposition of HNO₃ in spring. The estimation
246 in our study is 70% for EA and 61% for SA, close to Bey's result. Li et al. (2004) reported that
247 about 20% of anthropogenic NO_x emitted by NA is deposited out of its boundary (about 1000 km
248 offshore). Stohl et al. (2002) calculated that 9-22% of surface NO_x emissions is exported out of
249 the boundary layer of NA. Our estimation is about 30%, higher than Li's and Stohl's results. HTAP
250 I study by Sanderson et al. (2008) developed a SR relationship for NO_y deposition among NA, EU,
251 SA and EA. Their results showed that about 12-24% of the emitted NO_x is deposited out of source
252 regions. This study of HTAP II finds a higher percentage of export (26-34%). It should be noted
253 that the estimations of different studies are influenced by several factors and the results are not
254 fully comparable: (1) Definition of boundaries of source and receptor regions. For instance, Li et
255 al. (2004) defined the boundary of NA by a squared domain: 65-130°W, 25-55°N, while we use
256 the continental boundaries defined by HTAP II. There are also changes in the definition of
257 boundaries from HTAP I to HTAP II. For instance, HTAP I includes Mexico and Central America
258 in NA, but they are defined as a separate region in HTAP II (region 12 in Fig. 1). The boundary of
259 EU is also changed in HTAP II. (2) HTAP I simulations only change the NO_x emission, but HTAP
260 II simulations also reduce the other anthropogenic emissions, including SO₂ and PM. The joint
261 control of multiple species may result in more reduction in NO_y deposition and it is hard to estimate
262 this effect in this study.

263 **3.2 Impact of hemispheric transport on deposition**

264 This section investigates the impact of hemispheric transport on deposition in the receptor regions.
265 Figure 2 is annual response of S deposition to 20% emission reduction in source regions calculated
266 as Eq. (2).

$$267 \quad \text{Response} = \frac{\Delta \text{Depo} (\text{perturbation})}{\text{Depo} (\text{base})} \times 100\% \quad (2)$$

268 where $\Delta \text{Depo} (\text{perturbation})$ is the ΔDepo between perturbation case and base case. $\text{Depo} (\text{base})$
269 is the deposition under base case. The negative values mean that the deposition decreases with
270 reduction in emission. Table S2 summarizes the regional median deposition fluxes under base case
271 and under emission perturbation cases. Fig. 2(a) shows the global response of S deposition to 20%
272 emission reduction in NA. The largest deposition change is found in the source region NA, with a

273 4-20% decrease in S deposition in the non-coastal region and 14-16% decrease at the east coast.
274 The impact on the North Atlantic Ocean deposition declines gradually from near coastal region
275 (12-14%) to the open ocean (2-12%) and Eurasia (<1%). Fig. 2(b) shows the global response of S
276 deposition to 20% emission reduction in EU. Although the impact on continental non-coastal
277 regions is high (6-18%), the impact on the coastal regions is generally less than 6%, much lower
278 than NA's impact on its east coast (14-16%). The deposition in North Africa, central Asia and
279 western RU is affected by 2-6%. The 20% emission reduction in SA (Fig. 2(c)) shows large
280 influence over its south-west coast and the Arabian Sea (4-12%). The SA's outflow affects the
281 deposition in southeastern ME and eastern Sub Saharan Africa by 1-4% and western EA and
282 Southeast Asia (mainly Bangladesh) by 2-6%. Fig. 2(d) shows the impact on S deposition from
283 20% emission reduction on deposition in EA. On one hand, the impact is strong on the east coast
284 of China (12-16%) and decreases gradually over the North Pacific Ocean (4-10%). Although the
285 majority of S emission is deposited on the North Pacific Ocean, the influence on western NA can
286 still reach 4-6%. On the other hand, the impact on Southeast Asia and SA is much lower (2-5%
287 and <1%), due to the block of air flows by the Himalaya Mountains (Fig. S4). The 20% emission
288 reduction in ME mainly affects the S deposition in Africa by 2-10% and western EA by 2-4%. Fig.
289 2(f) shows the S deposition change with 20% emission reduction in RU. The regions of impact are
290 mainly at high latitudes in the North Hemisphere, including northern EU (2-6%) and western
291 Arctic Circle (1-4%). The Russian flow enters the Arctic in the lower troposphere in winter season
292 (Stohl, 2006).

293 The impact of NO_x emission reduction on NO_y deposition from each source region is
294 shown in Fig 3. The overall impact is qualitatively similar to that of S emission in the spatial
295 pattern, with some differences in the values. For some regions, the impact of hemispheric transport
296 on NO_y deposition is lower than that on S deposition. For instance, SA's impact on eastern Africa
297 is about 1-4% on S deposition, but is <1% on NO_y deposition. ME's impact on the western Africa
298 and Gulf of Guinea is about 2-4% on S deposition, but is <1% on NO_y deposition. These smaller
299 sensitivities reflect differences in lifetimes, and the lower formation of aerosol nitrate under warm
300 conditions in tropical regions. Under the NA perturbation case (Fig. 3(a)), an 8-12% change of
301 NO_y deposition is found on the west coast of California, due to high NO_x emission in California
302 from mobile source, which is not seen in S deposition. The impact of emission reduction in EU
303 and EA on their coastal regions is generally 2-4% higher for NO_y deposition than S deposition

304 (Fig. 3(b) and (d)). The impact on NH_x deposition is similar to that on NO_y deposition (Fig. S2).
305 It should be noted that this is the result from 20% emission reduction in the source regions,
306 therefore the actual impact (100% emission reduction) could be 5 times higher when assuming a
307 linear relationship between 20% and 100% emission reduction on deposition.

308 We quantify the amount of deposition carried by hemispheric transport and study its
309 seasonality. Fig. 4 shows the monthly changes of S deposition for 20% emission reductions in
310 source regions. The values are meridional sum with a west-east resolution of 0.1 degree, and
311 display well the locations of the source regions. The negative values indicate the amounts of
312 pollutants transported from source regions to receptor regions. According to Fig 4(a), NA has
313 about $(1-10) \times 10^4 \text{ kg(S) month}^{-1}$ per 0.1° longitude of its S emission transported and deposited
314 over the North Atlantic Ocean ($15-75^\circ\text{W}$) throughout the whole year. We also find about $(1-3)$
315 $\times 10^4 \text{ kg(S) month}^{-1}$ per 0.1° longitude decrease of S deposition at about 90°E and 120°E in spring
316 and fall, which gives evidence to NA's influence on Eurasia via transatlantic flow, although this
317 amount accounts for less than 1% of local S deposition (white in Eurasia in Fig.2(a)). Fig. 4(b)
318 shows that about $(1-3) \times 10^4 \text{ kg(S) month}^{-1}$ per 0.1° longitude of EU's emission is transported to
319 and deposited at $30-60^\circ\text{E}$ in RU throughout the whole year and at $100-120^\circ\text{E}$ in EA in spring and
320 fall. According to Fig. 4(c), SA exports its S emission to $30-60^\circ\text{E}$ in ME and eastern Africa in
321 early spring and to $90^\circ\text{E}-180^\circ$ in EA and North Pacific Ocean from late spring until fall. In
322 particular, the influence on EA can reach $(5-10) \times 10^4 \text{ kg(S) month}^{-1}$ per 0.1° longitude in mid-
323 spring. According to Fig. 4(d), EA's S emission is widely transported and deposited over the North
324 Pacific Ocean throughout the whole year. The Asian outflow arrives at the west coast of NA
325 ($\sim 130^\circ\text{W}$) in all seasons except summer, but only reaches far in western NA ($\sim 90^\circ\text{W}$) in spring
326 and brings about $1 \times 10^4 \text{ kg(S) month}^{-1}$ per 0.1° longitude of S deposition. The export to SA is only
327 found during the Asia winter monsoon. The monthly changes of NO_y deposition with perturbation
328 experiments are shown in Fig. 5. Compared to S deposition, the change in NO_y deposition by
329 hemispheric transport is generally smaller. For instance, the NA's impact on Eurasia is $(1-3) \times 10^4$
330 kg(S) month^{-1} per 0.1° longitude for S deposition, but is less than $0.5 \times 10^4 \text{ kg(N) month}^{-1}$ per 0.1°
331 longitude for NO_y deposition. The SA's impact on EA ($90-120^\circ\text{E}$) can reach $(5-10) \times 10^4 \text{ kg(S)}$
332 month^{-1} per 0.1° longitude for S deposition, but the amount is 4 times lower for NO_y deposition.
333 This result is in accordance with the S-R results in section 3.1 that more S emission is transported
334 out of the source regions than N emission, probably due to longer chemical lifetimes and higher

335 emission altitudes. Patterns similar to NO_y are also found in the monthly changes of NH_x
336 deposition (Fig. S3).

337 The deposition change via transport between neighboring regions is found throughout the
338 whole year and is slightly stronger in winter, such as between EU and RU ($\sim 30^\circ\text{E}$) (Fig. 4(b) and
339 (f)) and from EA to the North Pacific Ocean ($\sim 130^\circ\text{E}$) (Fig. 4(d)). This is consistent with the
340 seasonality we found for the export of emission by source regions in section 3.1. In addition, most
341 source regions reduce more S and NO_x emissions in winter than the other seasons (Table S3), thus
342 more emissions are exported abroad in winter. On the contrary, the deposition change by transport
343 over long distance mainly occurs in spring and fall, especially for the hemispheric transport from
344 NA to EU, from EU to EA and from EA to NA. The seasonality of long-range transport for NA,
345 EU and EA well fits the characteristic of westerlies, which is the prevailing winds in the mid-
346 latitude of the North Hemisphere. This agrees with the seasonality of the transpacific, transatlantic
347 and trans-Eurasia flows of air pollutants that spring is the most efficient season for long-range
348 transport for mid-latitudes (Holzer et al., 2005; Liu et al., 2005; Liang et al., 2004; Brown-Steiner
349 and Hess, 2011; Li et al., 2014; Auvray and Bey, 2005; Wild et al., 2004; Liu et al., 2003). Although
350 the westerlies is also strong in winter, the pollution in the air flow is low, because the formation
351 of secondary species like PAN is suppressed by slow oxidation in cold environment (Berntsen et
352 al., 1999; Deolal et al., 2013; Moxim et al., 1996), which plays an important role as a reservoir for
353 NO_x in the long-range transport of air plumes (Lin et al., 2010; Hudman et al., 2004).

354 **3.3 Own region and foreign contributions on deposition**

355 This section compares the contributions between hemispheric transport and own region emission
356 control on deposition. A metric called extra-regional emission reduction (RERER) is calculated
357 by dividing the ΔDepo due to foreign emission reduction by ΔDepo due to global (foreign + own
358 region) emission control following Eq. (3).

359

$$360 \quad RERER_i = \frac{\Delta \text{Depo}_i(\text{foreign})}{\Delta \text{Depo}_i(\text{global})} \quad (3)$$

361 where i is the region of focus. $\Delta \text{Depo}_i(\text{foreign})$ is the ΔDepo in region i due to 20% foreign
362 emission reduction. It is calculated by subtracting the ΔDepo due to 20% own region emission

363 control from Δ Depo due to 20% global emission reduction. Δ Depo_i (global) is the Δ Depo in
364 region i due to 20% global emission reduction. This metric indicates the importance of foreign
365 emission on local deposition. A low RERER value (close to 0) indicates a predominance effect of
366 own region emission on local deposition, while high RERER value (close to 1) means strong
367 impact of hemispheric transport on local deposition.

368 Table 2 shows the RERER values for total (include both non-coastal and coastal regions),
369 non-coastal and coastal regions. For both non-coastal and coastal regions, the own region impact
370 includes control of both its coastal and non-coastal regions, and the foreign impact comes from
371 emission reduction of foreign coastal and non-coastal regions. As we expected, NA (0.07, 0.17,
372 and 0.17), SA (0.04, 0.17, and 0.18) and EA (0.16 and 0.16) regions have relatively low RERER
373 values, due to large local emissions compared to the foreign contributions. EU (0.12, 0.34 and
374 0.36) and ME (0.32, 0.32 and 0.42) have relatively higher RERER values. RU is the only region
375 with RERER (0.55, 0.59 and 0.61) higher than 0.5, which means its deposition is almost equally
376 sensitive to foreign impact and own region control. The RERER values NO_y deposition are of
377 similar magnitudes to S deposition, while the RERER of NH_x deposition is 0.1 lower, probably
378 due to the lack of data from EA perturbation case, so that the contribution from EA is not
379 considered.

380 The RERER values of coastal regions are generally 0.1-0.3 higher than those of non-coastal
381 regions. The values for RU's coast are all higher than 0.84. Even regions with low non-coastal
382 RERER such as NA and SA have high RERER on coastal regions. For instance, the RERER of
383 NA reaches 0.3-0.4 for its coastal region, more than double of the RERER on its non-coastal
384 regions (0.05-0.12). Coastal regions receive high proportion of deposition from foreign transport.
385 Except large scale circulation like prevailing westerlies, the coastal regions are featured with
386 complex small scale circulations. For instance, the low-level jet (zonal winds with high speed)
387 contributes to the rainfall in coastal regions in Asia (Xavier et al., 2018). The orographic effects
388 enhance the precipitation over coastal mountain regions such as west coast of NA, EU and
389 southeast coast of RU (James and Houze, 2005). According to table 1, EA exports 5% of its S and
390 N emission to RU, almost half of which is deposited on RU's coastal regions. RU exports 7-12%
391 of S and N emission to EU, of which 30% is deposited on EU's coastal regions. The impact of
392 hemispheric transport is identical or even larger than the effect of own region emission control for

393 some coastal or near coastal regions. According to Fig. 2, 20% emission reduction in EA can
394 reduce 2-6% of S deposition in the west coast of NA. This effect is even larger than 20% emission
395 reduction in own region emission (<1%). Similarly, 20% emission reduction in NA can change 2-
396 5% of S deposition in west coast of EU, which is almost identical to the effect of 20% emission
397 control in EU. On one hand, the emissions in western NA and western EU are relatively low, thus
398 the effect of own region control is not significant. On the other hand, these coastal regions are in
399 the downwind location of eastern EA and eastern NA, which are the main source regions of S and
400 N emissions. Coastal regions serve as transit places for air-sea exchange with vulnerable
401 ecosystem (Jickells, 2006;Jickells et al., 2017). The over-richness of deposition in coastal water
402 and ecosystem can evoke a number of environmental issues, of which some are specifically for
403 coastal regions such as threats to coastal benthic and planktonic system and sustainability of
404 fishery (Paerl, 2002;Doney et al., 2007).

405 Figure 6 compares the percentage concentrations on deposition between foreign transport
406 and own region emission. Other (OTH, pattern fill in the figure) is calculated as $\Delta \text{Depo}_{(\text{GLO})} - \sum$
407 $\Delta \text{Depo}_{(\text{case})}$ (case = 6, including NA, EU, SA, EA, ME and RU). It indicates the deposition change
408 due to other reasons than the total effects of separate emission reduction in the 6 regions. It could
409 come from the emission reduction in rest of world, especially nearby regions such as from Central
410 Asia and North Africa to EU and ME. It could also come from the joint effects of emission control
411 by multiple source regions, which possibly change the oxidant chemistry, atmospheric mixing and
412 lifetimes of reactive pollutants. However, the model simulations do not allow to separate these two
413 contributions in this study. For regions with low RERER values (NA, SA and EA), the own region
414 emission dominates the deposition by more than 80%. For these regions, determined by vicinity
415 and transport patterns, the foreign impact is somewhat dominated by certain source regions, such
416 as from EA to NA (2-4% out of 4-5%), from ME to SA (5-6% out of 7-11%) and from SA to EA
417 (3-4% out of 4-7%). For EU and ME, there is about 20% contribution from “OTH”. Beside this,
418 RU contributes 4-5% to EU’s deposition and EU contributes 5% to ME’s deposition. For high
419 RERER regions, RU has a different pattern than the other regions. The contributions of
420 hemispheric transport from the other 5 regions (23-45%) are almost equivalent to its own region
421 emission control (39-45%), with significant contributions from EA (20-24%) and EU (13-15%).

422 Figure 7 shows the inter-model variation on simulating the ΔDepo of S, NO_y and NH_x
423 under emission perturbation cases, separately for wet and dry deposition. The values are global

424 integrated changes in component deposition for perturbation experiments from MMM results, with
425 error bars showing the maximum and minimum values of models. The figure only shows main
426 compositions of S and N deposition, which together account for more than 95% of total deposition.
427 In terms of S deposition (Fig. 7(a)), the modeled Δ Depo by multiple models, defined as (maximum
428 value of multi-model – minimum value of multi-model), ranges (0.06-0.23) Tg(N) yr⁻¹ and (0.01-
429 0.22) Tg(N) yr⁻¹ for SO₂ dry and wet deposition, and (0.01-0.03) Tg(N) yr⁻¹ and (0.009-0.17) Tg(N)
430 yr⁻¹ for SO₄²⁻ dry and wet deposition, respectively. High uncertainty is found in EA perturbation
431 case, where the model divergence are mainly from SO₂ wet and dry deposition and SO₄²⁻ wet
432 deposition. In terms of NO_y deposition (Fig. 5(b)), the differences among models range (0.003-
433 0.07) Tg(N) yr⁻¹ for NO₂ dry deposition, and (0.07-0.55) Tg(N) yr⁻¹ and (0.03-0.75) Tg(N) yr⁻¹ for
434 NO₃⁻ dry and wet deposition, respectively. The EA perturbation case also has the largest inter-
435 model variation, with high uncertainty in simulating both the NO₃⁻ wet and dry deposition. In terms
436 of NH_x deposition (Fig. 5(c)), the differences among models range (0.04-0.09) Tg(N) yr⁻¹ for NH₃
437 dry deposition, and (0.008-0.15) Tg(N) yr⁻¹ and (0.002-0.11) Tg(N) yr⁻¹ for NH₄⁺ dry and wet
438 deposition, respectively. Both EA and SA perturbation cases have relatively high uncertainties on
439 NH₄⁺ dry deposition. Overall, the inter-model variation is considerably high under emission
440 perturbation in Asian regions, especially EA. On one hand, the EA perturbation case assumes the
441 largest amount of emission reductions among all cases (Table S3). On the other hand, model
442 evaluation (Tan et al., 2018) reported high model bias in simulating the deposition in this region,
443 and suggest an incomplete knowledge from the combined picture provided by observations and
444 models.

445 **4 Conclusion**

446 This study assesses the impact of hemispheric transport on S and N deposition for 6 regions: North
447 America (NA), Europe (EU), South Asia (SA), East Asia (EA), Russia (RU) and Middle East
448 (ME), by using multi-model ensemble results from 11 models of HTAP II, with simulations under
449 base case and 20% emission perturbation scenario for each region.

450 We investigate the export of S and N emissions from source regions. Results show that
451 about 27-58%, 26-46% and 12-23% of the emitted S, NO_x and NH₃ emissions are deposited outside
452 of the source regions (ranges are for different source regions). The most significant exports of
453 emissions are: (1) transport between EU and RU. 10-14% of EU's emission is transported to RU

454 and 7-12% of RU's emission is transport to EU. (2) transport between EA and RU. 5% of EA's
455 emission is transported to RU and 4-5% of RU's emission is transported to EA. (3) transport from
456 SA to EA (4-9%). Most regions export 5-10% more emission in winter than summer, which is
457 highly influenced by chemistry, precipitation amount and frequency, atmospheric mixing and
458 transport patterns.

459 We explore the impact of hemispheric transport on deposition in receptor regions. Overall,
460 20% emission reduction in source regions could affect 1-10% of deposition on foreign continental
461 regions and 1-14% on foreign coastal regions and the open ocean. The most significant impacts
462 are from NA to the North Atlantic Ocean (2-14%), and from EA to North Pacific Ocean (2-12%)
463 and to western NA (4-6%). The amounts of deposition brought by hemispheric transport range
464 from 10^4 - 10^5 kg(S or N) month⁻¹ per 0.1° longitude (meridional sum). The impact on deposition
465 via short-distance transport between neighbouring regions (i.e. EU to RU) is generally found
466 throughout the whole year and slightly stronger in winter, while the long-range transport (i.e. from
467 EA to NA) mainly occurs in spring and fall.

468 We compare the impact from own region emission and foreign transport on local
469 deposition. The deposition in NA, SA and EA is dominated (~80%) by their own region emission,
470 while EU, ME and RU receive 40-60% of deposition from hemispheric transport. In particular,
471 Russian deposition is even equally contributed by foreign inputs and own region emission, with
472 high contributions from two neighbouring source regions: EA (~20%) and EU (~15%). For some
473 regions, upmost half of the hemispheric transport from foreign regions is deposited over their
474 coastal regions. Deposition in coastal regions or the near-coastal open ocean is found twice more
475 sensitive to long-range transport than non-coastal regions. For some coastal regions such as west
476 coast of NA and west coast of EU, the impact of hemispheric transport is identical or even larger
477 than that of own region emission control.

478 This study highlights the impact of hemispheric transport on deposition in coastal regions
479 and the open ocean, which hasn't been fully studied in the literature. We therefore suggest further
480 research on this impact on the mitigation of coastal and oceanic ecosystem, with regards to the
481 increasing air pollutants in hemispheric outflow. We also find significant impact of hemispheric
482 transport on deposition in relatively low emission regions such as RU. The impact on their
483 ecosystem and human health requires further research. Meanwhile, there is still a portion of foreign
484 impact that hasn't been attributed in this study (aggregated as other regions "OTH" in Fig. 6). For

485 instance, at least 4 regions (NA, EU, SA and ME) have shown considerable impact (2-10%) on
486 the S and N deposition in North Africa. But since North Africa is not included as a receptor/source
487 region in the perturbation experiments, it is hard to quantify the impact of long-range transport on
488 its deposition. Meanwhile, Southeast Asia is regarded as a big emission contributor in Asia. It is
489 important to establish an S-R relationship with other Asian regions. We suggest the future HTAP
490 simulations to include these regions in the perturbation experiments.

491

492 *Data availability.* The model data can be downloaded from AeroCom database
493 (<http://iek8wikis.iek.fz-juelich.de/HTAPWiki/FrontPage>) upon request.

494

495 *Competing interests.* The author declare that they have no conflict of interest.

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807 Figure captions:

808 Fig. 1. Administration boundaries of regions and coastal areas. 6 Regions with perturbation
809 experiments: 3-North America (NA), 4-Europe (EU), 5-South Asia (SA), 6-East Asia (EA), 11-
810 Middle East (MD) and 14-Russia, Belarussia, Ukraine (RU). Other regions: 1-Global, 2-Ocean
811 (including Arctic), 7-Southeast Asia, 8-Australia, 9-North Africa, 10- Sub Saharan Africa, 12-
812 Mexico, Central America, Caribbean, Guyanas, Venezuela, Columbia (Central America), 13-
813 South America, 15-Central Asia and 17-Antarctic.

814 Fig. 2 The response of S deposition to 20% emission reduction in source regions. The values are
815 the percentage changes (%) in deposition calculated as (changes in deposition with 20% emission
816 reduction) / (base case deposition) $\times 100\%$. The unit is % per $0.1 \times 0.1^\circ$ grid box.

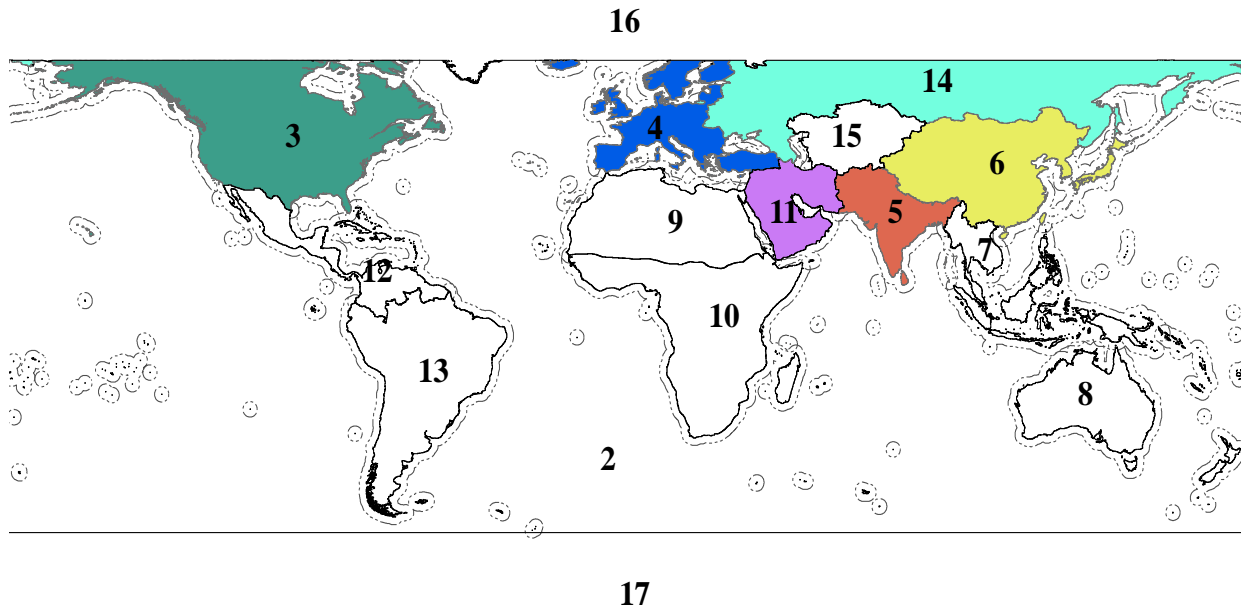
817 Fig.3 Same as Fig.2 but for NO_y deposition. The unit is % per $0.1 \times 0.1^\circ$ grid box.

818 Fig. 4 The monthly changes of S deposition with 20% emission reduction in source regions. The
819 values are meridional total values versus time with a west-east resolution of 0.1 degree. The unit
820 is $\times 10^4 \text{ kg(S) month}^{-1}$ per 0.1° longitude. The negative values indicate decline in deposition with
821 reduction in emission.

822 Fig. 5 Same as Fig.4 but for NO_y deposition. The unit is $\times 10^4 \text{ kg(N) month}^{-1}$ per 0.1° longitude.

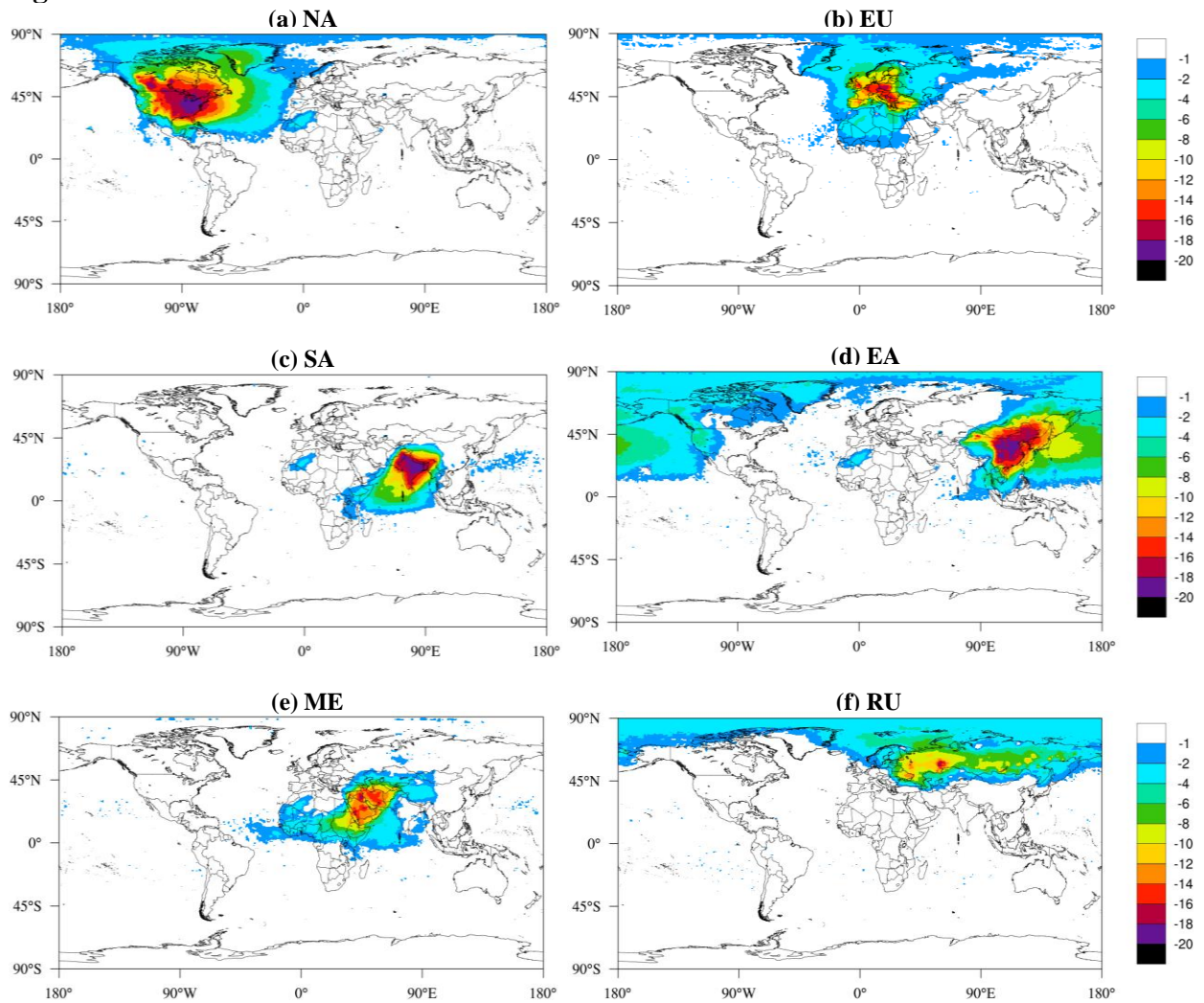
823 Fig. 6 Own region and foreign contributions on own region deposition. The values are calculated
824 by changes with 20% emission reduction. Other (OTH, pattern fill) is the contribution by other
825 reasons than emission reduction in the 6 regions (see text for details).

826 Fig. 7 Inter-model variations in deposition changes (unit: $\text{Tg(S or N) yr}^{-1}$) under emission
827 perturbation experiments. The values are MMM with error bars showing the max and min values
828 among all models.
829



831
832 Fig. 1. Boundaries of regions and coastal areas (dashed). 6 Regions with perturbation experiments:
833 3-North America (NA), 4-Europe (EU), 5-South Asia (SA), 6-East Asia (EA), 11-Middle East
834 (ME) and 14-Russia, Belarussia, Ukraine (RU). Others: 1-Global, 2-Ocean (including Arctic), 7-
835 Southeast Asia, 8-Australia, 9-North Africa, 10- Sub Saharan Africa, 12- Mexico, Central America,
836 Caribbean, Guyanas, Venezuela, Columbia (Central America), 13-South America, 15-Central Asia,
837 16-Arctic and 17-Antarctic.
838

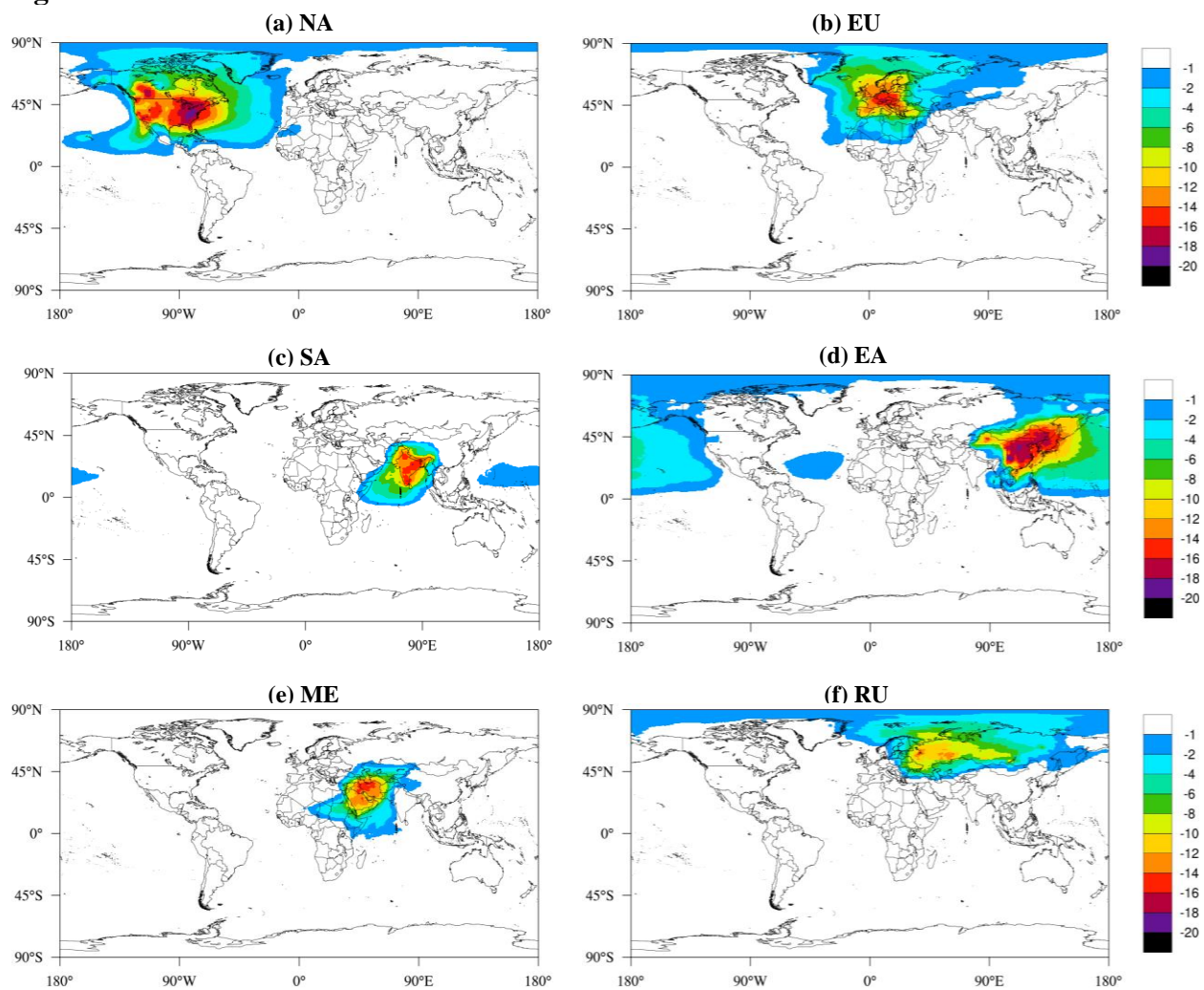
839 **Fig. 2**



840
841 Fig. 2 The response of S deposition to 20% emission reduction in source regions. The values are
842 the percentage changes (%) in deposition calculated as (changes in deposition with 20%
843 emission reduction) / (base case deposition) $\times 100\%$. The unit is % per $0.1 \times 0.1^\circ$ grid box.
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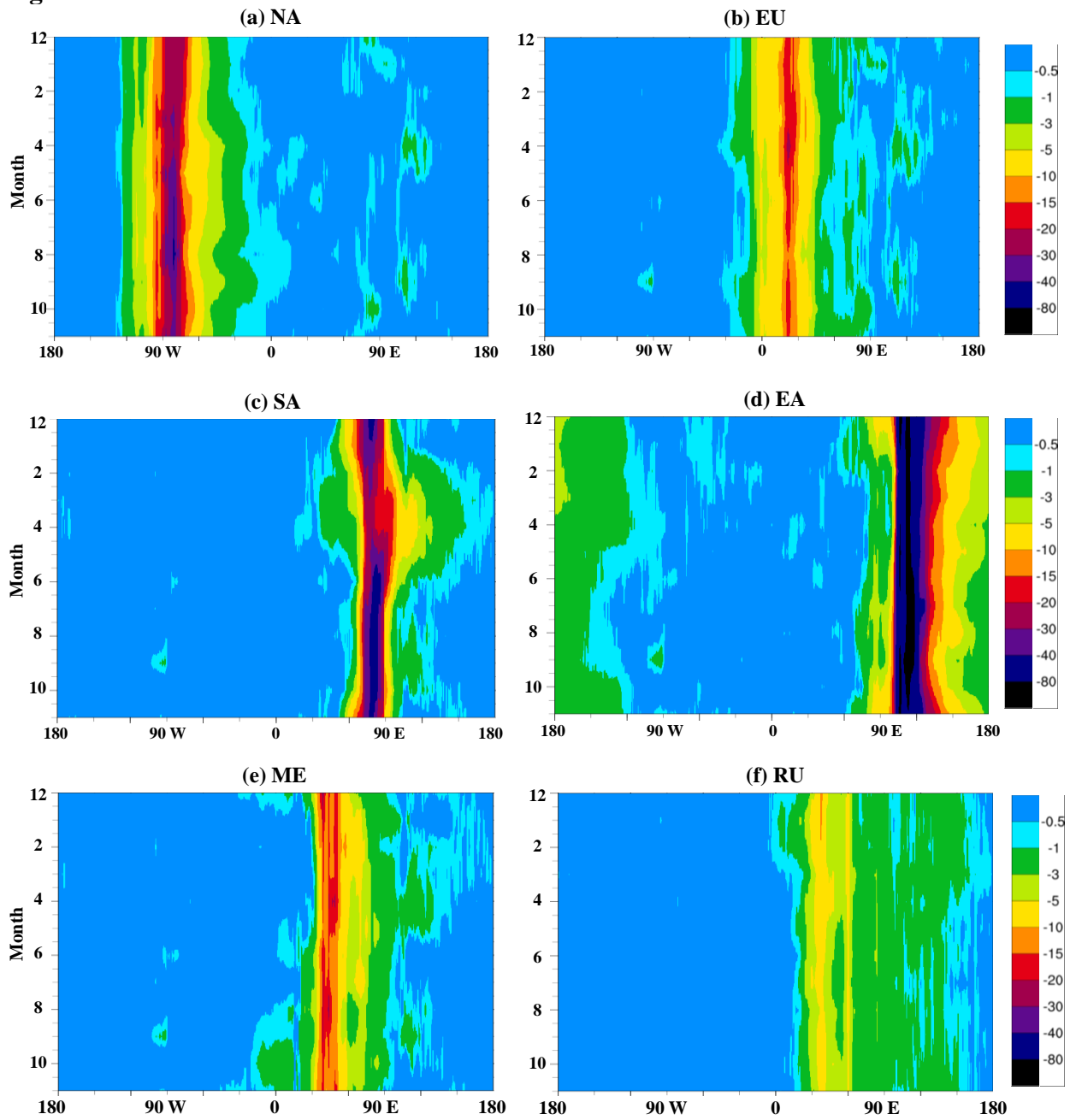
846 **Fig. 3**



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Fig.3 Same as Fig.2 but for NO_y deposition. The unit is % per $0.1 \times 0.1^\circ$ grid box.

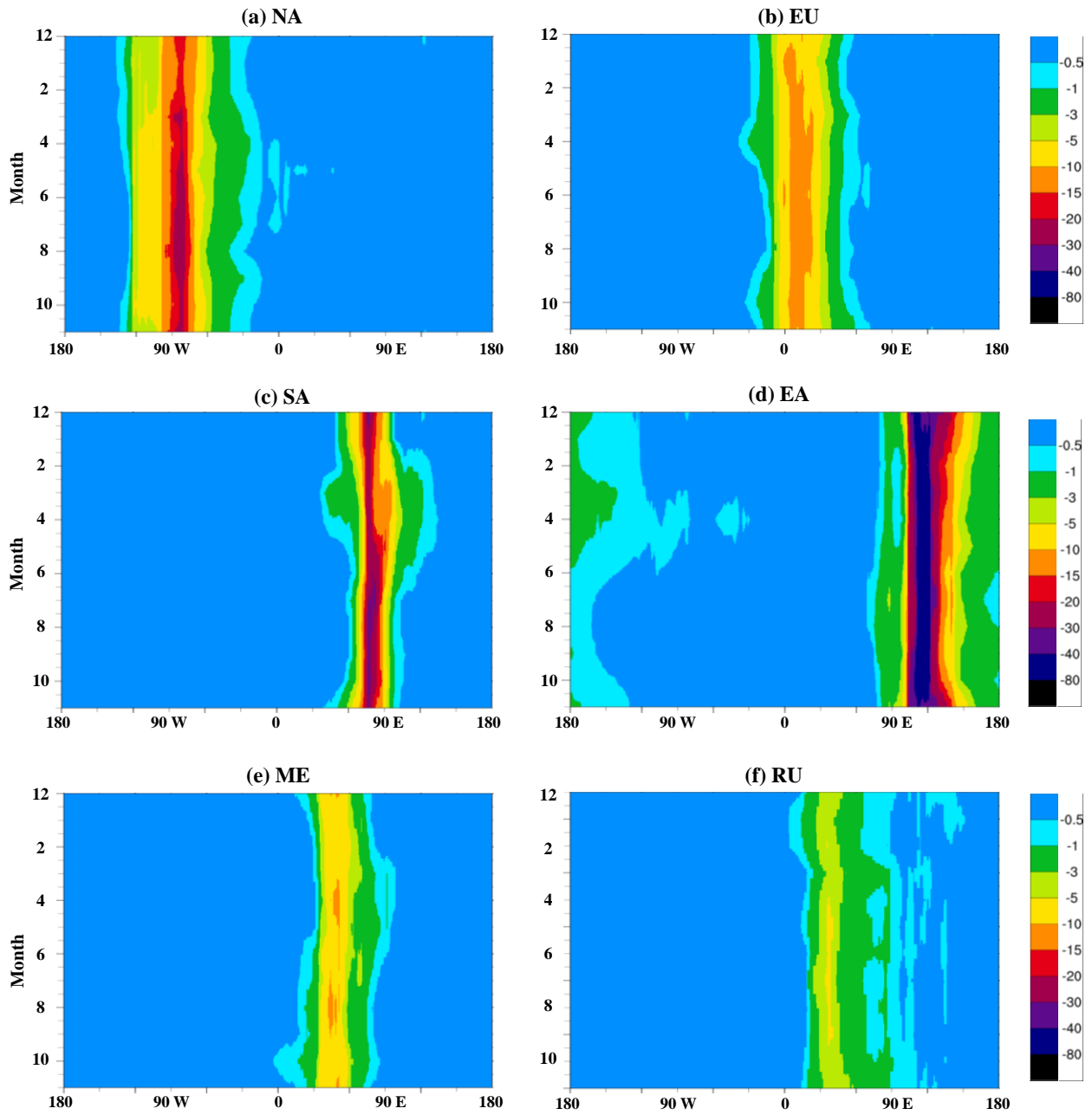
850 **Fig. 4**



851

853 Fig. 4 The monthly changes of S deposition with 20% emission reduction in source regions. The
854 x-axis values are meridional total values versus time (y-axis) with a west-east resolution of 0.1
855 degree. The unit is $\times 10^4$ kg(S) month⁻¹ per 0.1° longitude. Negative values indicate decline in
856 deposition with reduction in emission.

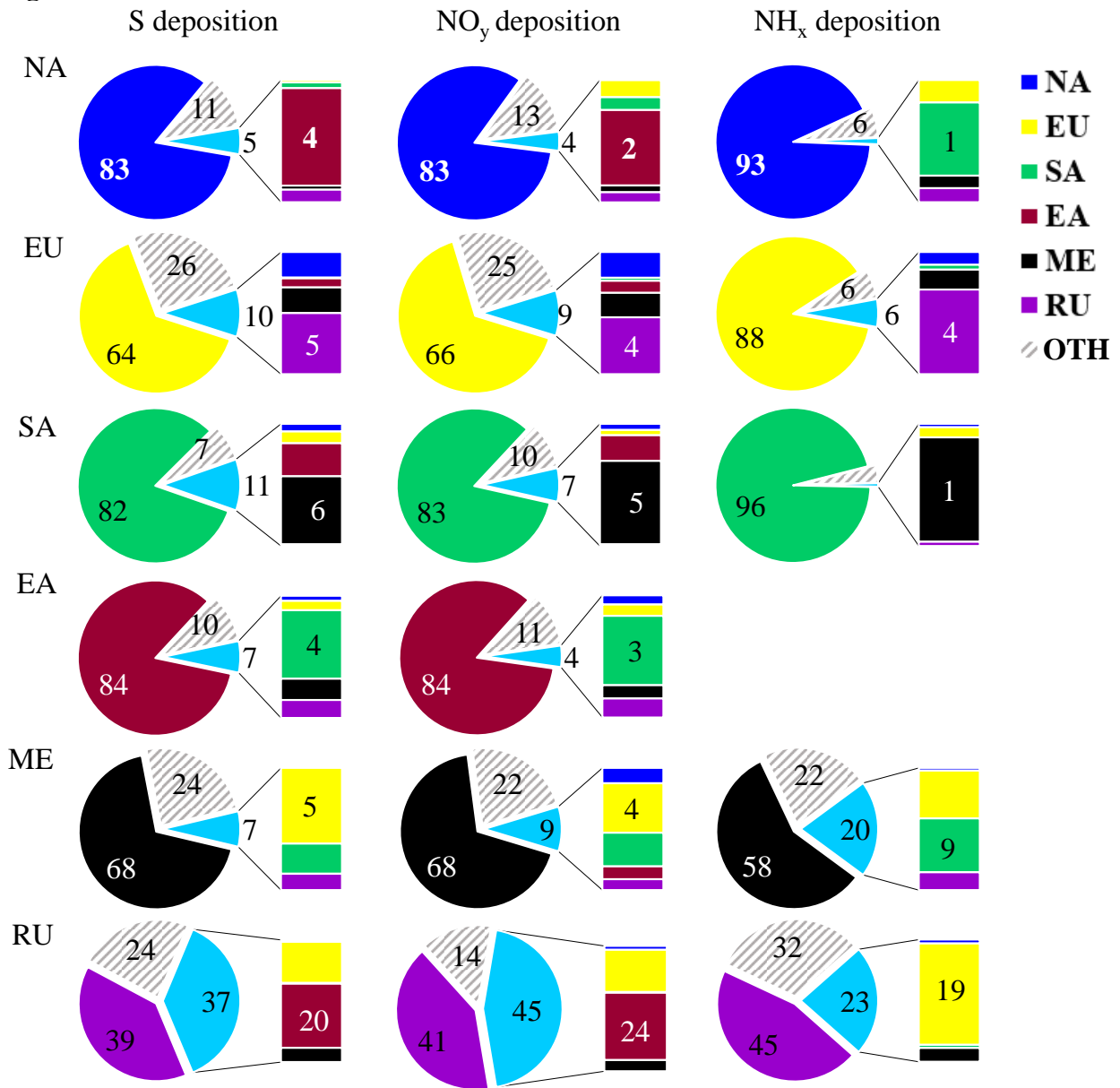
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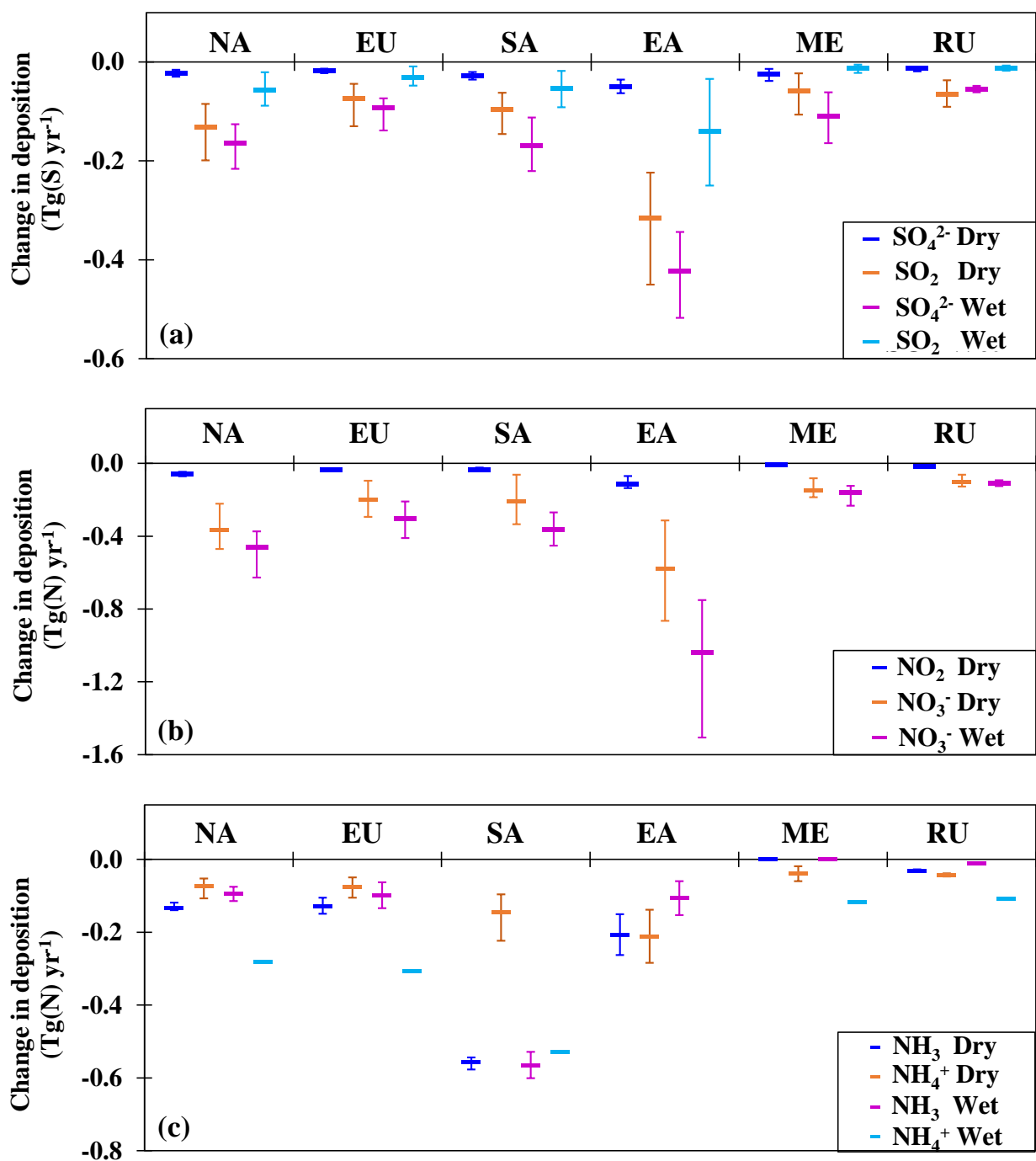
Fig. 5 Same as Fig.4 but for NO_y deposition. The unit is $\times 10^4 \text{ kg(N) month}^{-1}$ per 0.1° longitude.

862 **Fig. 6**



863
 864 Fig. 6 Own region and foreign contributions on own region deposition. The values are calculated
 865 by changes with 20% emission reduction. Other (OTH, pattern fill) is the contribution by other
 866 reasons than emission reduction in the 6 regions (see text for details).

867



869
 870 Fig. 7 Inter-model variations in wet and dry deposition changes (unit: Tg(S or N) yr⁻¹) under
 871 emission perturbation experiments. The values are global integrated changes in components of S
 872 and N deposition for perturbation experiments from MMM results with error bars showing the
 873 max and min values among all models. Species without error bars are derived from results of a
 874 single model.

876 Tables

877 Table 1. Source-receptor relationship of S/NO_y/NH_x deposition (%) for regions (including
 878 continental coastal and non-coastal regions). The values in the parentheses are for coastal regions
 879 as a subset of the total.

| | Receptor Regions | Source Regions | | | | | |
|-------------------------------|---------------------|----------------|-------------|-------------|-------------|------------|------------|
| | | NA | EU | SA | EA | ME | RU |
| S Deposition | NA | 68.9 (8.9) | 0.2 (0.1) | 0.2 (0.0) | 1.5 (0.6) | 0.3 (0.1) | 1.2 (0.6) |
| | EU | 1.1 (0.6) | 60.4 (14.4) | 0.0 (0.0) | 0.2 (0.1) | 2.1 (0.2) | 6.9 (2.5) |
| | SA | 0.5 (0.1) | 1.2 (0.3) | 66.4 (10.0) | 0.9 (0.4) | 7.9 (1.6) | 0.3 (0.1) |
| | EA | 0.6 (0.2) | 1.8 (0.4) | 8.8 (1.3) | 73.4 (11.5) | 4.6 (0.8) | 5.2 (1.4) |
| | ME | 0.0 (0.0) | 2.6 (0.6) | 0.6 (0.3) | 0.0 (0.0) | 42.4 (8.2) | 0.8 (0.2) |
| | RU | 0.4 (0.1) | 13.6 (2.2) | 0.1 (0.1) | 5.1 (2.2) | 5.0 (1.1) | 62.2 (4.4) |
| | Others | 28.5 | 20.1 | 23.8 | 19.1 | 37.7 | 23.4 |
| NO _y Deposition | NA | 71.5 (7.8) | 0.8 (0.2) | 0.5 (0.1) | 1.0 (0.3) | 0.5 (0.1) | 1.0 (0.3) |
| | EU | 1.3 (0.6) | 66.2 (17.5) | 0.2 (0.1) | 0.3 (0.1) | 3.5 (0.9) | 9.8 (2.9) |
| | SA | 0.2 (0.0) | 0.2 (0.0) | 66.2 (8.6) | 0.5 (0.2) | 7.9 (1.3) | 0.2 (0.0) |
| | EA | 0.6 (0.1) | 1.2 (0.2) | 6.2 (0.7) | 74.4 (14.3) | 2.4 (0.3) | 4.3 (0.9) |
| | ME | 0.4 (0.1) | 1.6 (0.3) | 0.9 (0.4) | 0.1 (0.0) | 54.4 (8.0) | 0.8 (0.2) |
| | RU | 0.6 (0.1) | 10.3 (1.3) | 0.1 (0.0) | 5.1 (2.2) | 4.9 (1.3) | 61.4 (3.1) |
| | Others | 25.6 | 19.7 | 25.8 | 18.6 | 26.4 | 22.5 |
| NH _x Deposition | NA | 88.4 (5.6) | 0.2 (0.1) | 0.3 (0.1) | -* | 0.7 (0.3) | 0.4 (0.2) |
| | EU | 0.6 (0.3) | 83.2 (17.8) | 0.0 (0.0) | - | 4.6 (1.2) | 11.9 (3.1) |
| | SA | 0.0 (0.0) | 0.1 (0.0) | 85.1 (7.6) | - | 8.6 (2.4) | 0.0 (0.0) |
| | EA | 0.0 (0.0) | 0.4 (0.1) | 4.2 (0.3) | - | 2.6 (0.5) | 3.8 (1.0) |
| | ME | 0.1 (0.0) | 1.3 (0.3) | 0.4 (0.2) | - | 49.4 (5.9) | 1.5 (0.4) |
| | RU | 0.4 (0.1) | 10.3 (1.3) | 0.1 (0.0) | - | 7.3 (1.5) | 76.9 (4.1) |
| | Others | 10.5 | 4.4 | 9.7 | - | 26.9 | 5.7 |

880 * Lack of NH₄⁺ wet deposition under EA emission perturbation experiment from all models.

881

882 Table 2. Extra-regional emission reduction (RERER) values of S/NO_y/NH_x deposition for
 883 continent non-coastal and coastal regions. The RERER values are calculated by dividing the Δ
 884 Depo due to foreign emission reduction by Δ Depo due to global (foreign + own region)
 885 emission control. Total column gives the RERER for coastal and non-coastal together.

| Regions | S deposition | | | NO _y deposition | | | NH _x deposition | | |
|---------|--------------|-------------|---------|----------------------------|-------------|---------|----------------------------|-------------|---------|
| | Total | Non-coastal | Coastal | Total | Non-coastal | Coastal | Total | Non-coastal | Coastal |
| NA | 0.17 | 0.12 | 0.40 | 0.17 | 0.12 | 0.43 | 0.07 | 0.05 | 0.31 |
| EU | 0.36 | 0.27 | 0.53 | 0.34 | 0.27 | 0.48 | 0.12 | 0.09 | 0.22 |
| SA | 0.18 | 0.14 | 0.35 | 0.17 | 0.12 | 0.37 | 0.04 | 0.03 | 0.17 |
| EA | 0.16 | 0.14 | 0.28 | 0.16 | 0.12 | 0.27 | - | - | - |
| ME | 0.32 | 0.27 | 0.46 | 0.32 | 0.27 | 0.50 | 0.42 | 0.36 | 0.67 |
| RU | 0.61 | 0.56 | 0.84 | 0.59 | 0.52 | 0.90 | 0.55 | 0.49 | 0.85 |

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