

1 **Single-particle characterization of aerosols collected at a remote site in the Amazonian**
2 **rainforest and an urban site in Manaus, Brazil**

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17
18 **Abstract**

19 In this study, aerosol samples collected at a remote site in the Amazonian rainforest and an urban
20 site in Manaus, Brazil, were investigated on a single particle basis using a quantitative energy-
21 dispersive electron probe X-ray microanalysis (ED-EPMA). Twenty-three aerosol samples were
22 collected in four size ranges (0.25-0.5, 0.5-1.0, 1.0-2.0, and 2.0-4.0 μm) during the wet season in
23 2012 at two Amazon basin sites: 10 samples in Manaus, an urban area; and 13 samples at an 80-m
24 high tower, located at the Amazon Tall Tower Observatory (ATTO) site in the middle of the
25 rainforest, 150 km northeast of Manaus. The aerosol particles were classified into nine particle
26 types based on the morphology on the secondary electron images (SEIs) together with the
27 elemental concentrations of 3,162 individual particles: (i) secondary organic aerosols (SOA), (ii)

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28 ammonium sulfate (AS), (iii) SOA and AS mixtures, (iv) aged mineral dust, (v) reacted sea-salts,
29 (vi) primary biological aerosol (PBA), (vii) carbon-rich or elemental carbon (EC) particles, such
30 as soot, tarball, and char, (viii) fly ash, and (ix) heavy metal (HM, such as Fe, Zn, Ni, and Ti)-
31 containing particles. In submicron aerosols collected at the ATTO site, SOA and AS mixture
32 particles were predominant (50-94% in relative abundance) with SOA and ammonium sulfate
33 comprising 73-100%. In supermicron aerosols at the ATTO site, aged mineral dust and sea-salts
34 (37-70%) as well as SOA and ammonium sulfate (28-58%) were abundant. PBAs were observed
35 abundantly in the PM_{2.4} fraction (46%), and EC and fly ash particles were absent in all size
36 fractions. The analysis of a bulk PM_{0.25-0.5} aerosol sample from the ATTO site using Raman
37 microspectrometry and attenuated total reflection Fourier transform infrared spectroscopy showed
38 that ammonium sulfate, organics, and minerals are the major chemical species, which is consistent
39 with the ED-EPMA results. In the submicron aerosols collected in Manaus, either SOA and
40 ammonium sulfate (17-80%) or EC particles (6-78%) were dominant depending on the samples.
41 In contrast, aged mineral dust, reacted sea-salt, PBA, SOA, ammonium sulfate, and EC particles
42 comprised most of the supermicron aerosols collected in Manaus. The SOA, ammonium sulfate,
43 and PBAs were mostly of a biogenic origin from the rainforest, whereas the EC and HM-containing
44 particles were of an anthropogenic origin. Based on the different contents of SOA, ammonium
45 sulfate, and EC particles among the samples collected in Manaus, a considerable influence of the
46 rainforest over the city was observed. Aged mineral dust and reacted sea-salt particles, including
47 mineral dust mixed with sea-salts probably during long-range transatlantic transport, were
48 abundant in the supermicron fractions at both sites. Among the aged mineral dust and reacted sea-
49 salt particles, sulfate-containing ones outnumbered those containing nitrates and sulfate+nitrate in
50 the ATTO samples. In contrast, particles containing sulfate+nitrate were comparable in number to
51 particles containing sulfate only in the Manaus samples, indicating the different sources and
52 formation mechanisms of secondary aerosols, i.e., the predominant presence of sulfate at the ATTO
53 site from mostly biogenic emissions and the elevated influences of nitrates from anthropogenic
54 activities at the Manaus site.

55

56 **1. Introduction**

57 The Amazonian rainforest is regarded as one of the primitive continental regions and
58 atmospheric aerosol particles over the region are expected to be influenced minimally by

59 anthropogenic activities, particularly during the wet season (Andreae, 2007; Martin et al., 2010b;
60 Chen et al., 2015). The unique near-natural conditions during the wet season make it an ideal place
61 to understand the occurrence, nature, origin, and transport of aerosol particles, which can directly
62 scatter and absorb solar radiation and indirectly serve as cloud condensation nuclei (CCN) and/or
63 ice nuclei (IN), to better predict the additional anthropogenic effects on aerosol particles, and to
64 help determine their influences on the environment, climate, and human health (Artaxo et al.,
65 2013). The Amazon Tall Tower Observatory (ATTO) consists of several observatory towers built
66 in the middle of the Amazon rainforest for a continuous and detailed study of biota-atmosphere
67 interactions (Andreae et al., 2015). Manaus, the capital of Amazonas state, is a large city located
68 in the northern region of Brazil with more than 2 million inhabitants in an area of 11,401 km²
69 (IBGE, 2017). The city, which is surrounded by the largest tropical rainforest, has a large industrial
70 zone, a port area at the Rio Negro, an energy matrix based on fuel oil, diesel, and natural gas, and
71 a growing automotive fleet (Martin et al., 2010a, 2016). Consequently, the pollution plume from
72 Manaus can act as a laboratory for examining the perturbations in natural processes (Martin et al.,
73 2016). Only a few studies examined airborne particles over the Amazon rainforest and nearby
74 urban sites simultaneously (Fraund et al., 2017; Martin et al., 2016). Therefore, there is still little
75 information on the urban vs. ecosystem influences. Especially, single-particle characterization of
76 aerosols collected at a Manaus city center has been scarce.

77 Many studies have been performed on the aerosol characteristics in the Amazon basin, but
78 the formation and dynamic processes of Amazonian aerosols are not completely understood
79 (Andreae et al., 2015; Martin et al., 2010a, 2016; Fraund et al., 2017; Fan et al., 2018). The Amazon
80 Basin atmosphere is near-pristine during the wet season, whereas biomass burning prevails during
81 the dry season (Andreae et al., 2007; Pöschl et al., 2010; Artaxo et al., 2013; Pöhlker et al., 2018).
82 Based on a long-term study, it was reported that aerosol particles in the coarse fraction are
83 relatively constant in concentrations through the wet and dry seasons, whereas the aerosol particle
84 levels in the fine fraction differ due to the predominant influence of biomass burning during the
85 dry season (Artaxo et al., 2013; Moran-Zuloaga et al., 2018). Scanning electron microscopy/energy
86 dispersive X-ray spectrometry (SEM/EDX) studies categorized the Amazonian aerosols mainly as
87 secondary organic aerosol (SOA) particles, sulfates/chlorides, primary biological aerosol (PBA)
88 particles, mineral dust, sea salts (fresh and/or aged), and pyrogenic carbon particles within the
89 different size fractions (Krejci et al., 2005; Pöschl et al., 2010). Over the Amazonian rainforest,

90 SOA particles are mainly formed through the condensation of biogenic organic compounds onto
91 biogenic K-rich salt particles emitted from the forest and are predominant in the fine fraction,
92 which are important for CCN (Pöhlker et al., 2012) and can also affect the potential of mineral
93 particles when acting as an organic coating (Möhler et al., 2008). Under high relative humidity
94 conditions, nano- and micrometer SOA particles with a dominance of α -pinene and isoprene as
95 their precursors can remain in the liquid phase (Bateman et al., 2016), which further enhances the
96 formation of SOA as well as the oxygen-to-carbon (O/C) ratios. Hence, the study of this particle
97 type can help elucidate some of the atmospheric interactions (Lin et al., 2014). The atmosphere in
98 the Amazon Basin is also rich in PBA particles (Andreae, 2007; Artaxo et al., 1998, 2013; Martin
99 et al., 2010a). Their unique morphology and elemental compositions of major C and O with minor
100 S, K, P, Na, N, Cl, and/or Mg obtained by SEM/EDX are characteristic of individual PBAs like
101 fungal spores (China et al., 2016). PBA particles can contribute to CCN after being transported to
102 cloud formation altitudes by strong convection (Artaxo et al., 2013). At the ATTO site, single
103 particle analysis by a combination of scanning transmission X-ray microscopy/near edge X-ray
104 absorption fine structure spectroscopy (STXM/NEXAFS) and SEM/EDX highlighted the
105 dominance of biological particles and the abundance of biogenic SOA and the presence of C, N,
106 O, P and K are characteristic of aerosols at the area (Fraund et al., 2017). Based on an investigation
107 on particulate matter during the wet season, oxidized organic components were significantly
108 observed at Manaus sites (de Sá et al., 2018), where one third of the potential SOA would be of an
109 urban origin (Palm et al., 2018).

110 In this study, twenty-three aerosol samples collected at the ATTO site and at an urban site
111 in Manaus during the wet season in 2012 were examined on a single particle basis using a
112 quantitative energy-dispersive electron probe X-ray microanalysis (ED-EPMA), which provided
113 information on the morphology and chemical compositions of aerosols containing both light and
114 heavy elements. This paper presents the different characteristics of the aerosols collected at the
115 rainforest and in Manaus.

116

117 **2. Experimental section**

118 **2.1. Samples**

119 During the wet season in 2012, aerosol samples were collected at two sampling sites on the
120 Amazon basin, i.e., ATTO and a central area of Manaus. The ATTO site (S 02° 08.647' W 58°

121 59.992') is situated in the Uatumã Sustainable Development Reserve, approximately 150 km
122 northeast of Manaus (Fig. 1). This is a multidisciplinary research site of an international joint
123 project between Brazil and Germany for continuous monitoring of the biological, physical, and
124 chemical functions of the Amazon rainforest to answer questions related to climate change
125 (Andreae et al., 2015). Aerosol sampling was performed at an 80-m-height walk-up tower at the
126 ATTO site. In Manaus, the sampling site is situated in the central part of the city (S 03° 05.753',
127 W 59° 59.419'), which is at a representative urban region influenced by electricity production
128 based on fuel oil, diesel, and natural gas, biogenic emissions from the surrounding forest, and
129 mostly by light duty (using gasoline and ethanol) vehicle traffic. Heavy vehicles that use diesel
130 account to less than 10 % of the urban fleet (Medeiros et al., 2017). The location is nearby a small
131 parking area and around 200 m away from the intersection of four busy avenues, with frequent
132 diurnal traffic jam on weekdays. Particles were collected at a 2 m height above ground level. The
133 aerosol samples were collected on TEM grids (Ted Pella Inc., USA, Ted Pella Inc.,
134 Carbon/Formvar 200 mesh Cu grid, 35–70 nm thickness) using a five stage Battelle impactor (the
135 cut-off diameters are 0.25, 0.5, 1, 2, and 4 μm for stages 1-5, respectively) at the ATTO and Manaus
136 sites on April 1, 16, 17, and 18 and May 1, 2, and 3 (the four- and three-day samples were collected
137 at the ATTO and Manaus sites, which are notated as samples SA1-SA4 and SM1-SM3,
138 respectively). On each date, the sampling started around noon (local time) and lasted for
139 approximately 100 min. The individual particles collected on stages 1-3 ($\text{PM}_{0.25-0.5}$, $\text{PM}_{0.5-1.0}$, and
140 $\text{PM}_{1.0-2.0}$) for each sample and on an additional stage 4 ($\text{PM}_{2.0-4.0}$) for the SA4 and SM3 samples
141 were examined.

142 During the sampling period, the temperature was in the range of 22 to 32°C and the relative
143 humidity was above 55%. On April 16 and May 1 and 2, rain events occurred within the previous
144 24 hours prior to sampling. Detailed information on sampling dates and times and the
145 meteorological conditions during the samplings are given in Table 1. The ten-day backward air-
146 mass trajectories were obtained using the Hybrid Lagrangian Single-Particle Integrated Trajectory
147 (HYSPLIT) model from the NOAA Air Resources Laboratory's web server
148 (<http://www.arl.noaa.gov/ready/hysplit4.html>), as shown in Fig. 2. In the HYSPLIT calculation,
149 meteorological data output from the Global Data Assimilation System (GDAS) using GDAS1 data
150 with a horizontal resolution of 1° corresponding to $\sim 100 \text{ km} \times 100 \text{ km}$ and 23 vertical layers was
151 used, which was reported to provide plausible backward trajectory analysis (Su et al., 2015). All

152 samples were influenced by transatlantic air masses at a 1000 m receptor height and the Manaus
153 site was influenced mainly by the surrounding rainforest at 500 m and 100 m heights.

154

155 **2.2. EPMA measurements and data analysis**

156 Low-Z particle EPMA was carried out by SEM (JSM-6390, JEOL) equipped with an
157 Oxford Link SATW ultrathin window EDX detector, which has a spectral resolution of 133 eV for
158 Mn K_{α} X-rays. The X-ray spectra were recorded using INCA Energy software. To achieve the
159 optimal experimental conditions, such as the low background level in the X-ray spectra and good
160 sensitivity for low-Z element analysis, an accelerating voltage of 10 kV, a beam current of 0.5 nA,
161 and a measuring time of 20 s were used. X-ray spectral data acquisition for individual particles
162 was carried out manually in point analysis mode, i.e., the electron beam was focused at the center
163 of each particle, and X-rays were acquired while the beam remained fixed on this single spot. The
164 secondary electron images (SEIs) and X-ray spectra of an overall 3,162 individual particles for the
165 ATTO and Manaus samples were examined. As the TEM grids are thin (35-70 nm thickness),
166 strong X-rays from the Al or Cu metal stub commonly used in the SEM/EDX measurement would
167 be a problem when the TEM grid substrate is placed on it. A home-made sample holder (Fig. 3(a))
168 for the TEM grid samples was used to avoid interference from the metal stub, resulting in X-ray
169 spectra of bare TEM grids, which showed only C and O X-ray peaks from their carbon/Formvar
170 thin-film, a Cu-L peak caused by lateral scattering from the Cu bars of TEM grids, and a Si peak
171 from an impurity, as shown in Fig. 3(b). The net X-ray intensities for the chemical elements were
172 obtained by non-linear, least-square fitting of the spectra using the AXIL program (Vekemans et
173 al., 1994). Although the characteristic X-ray intensities of C and O were low for the bare TEM
174 grids, determination of the C and O concentrations for individual particles on the TEM grids was
175 performed using a methodology based on the Monte Carlo calculation technique to correct for the
176 interfering X-ray peaks of C and O emitted from the TEM grid, which provided reliable
177 quantification results when applied to the quantification of standard (sub)micron particles, such as
178 CaCO_3 , CaSO_4 , Na_2SO_4 , and SiO_2 . On the other hand, electron beam-sensitive particles, such as
179 NaNO_3 , $\text{Ca}(\text{NO}_3)_2 \cdot 4\text{H}_2\text{O}$, and ammonium sulfate, provided deviating quantification results (Geng
180 et al., 2010). As the Cu-L and Si X-ray intensities from the bare TEM grids are quite small (< 20
181 cps) under these measurement conditions, the two peaks could be neglected safely during the
182 quantification procedure.

183

184 **3. Results and discussion**

185 **3.1. Particle types observed in samples collected at the ATTO and Manaus sites**

186 In this study, the analyzed particles were classified based on their X-ray spectral and SEI
187 data, where 9 different particle types were observed in the samples collected at the ATTO and
188 Manaus sites during the wet season in 2012; i.e., (i) SOA, (ii) ammonium sulfate (AS) particles,
189 (iii) SOA and AS mixture particles, (iv) aged (reacted) sea-salt, (v) aged mineral dust, (vi) PBA
190 particles, (vii) carbon-rich or elemental carbon (EC) particles such as soot, tarball, and char or coal
191 dust, (viii) fly ash particles, and (ix) heavy metal-containing (HM) particles. In a previous study
192 (Pöschl, et al., 2010), five types of aerosols, such as (i) SOA droplets, (ii) SOA-inorganic mixture
193 particles where the inorganics are mostly sulfates and/or chlorides, (iii) PBA, (iv) mineral dust,
194 and (v) pyrogenic carbon particles, were reported based on single particle analysis for aerosol
195 samples collected at a remote site north of Manaus, Brazil during the 2008 wet season (3-13 March)
196 using SEM/EDX. The pure SOA droplets dominated in the nucleation and Aitken modes, whereas
197 the pure SOA, SOA-inorganic mixture particles, and pyrogenic carbon particles dominated in
198 accumulation mode (Pöschl, et al., 2010). With the exception of the reacted sea-salt particles
199 probably from the Atlantic Ocean as well as the abundant ammonium sulfate aerosols, the particle
200 types observed in this study are comparable to their study. Figs. 4 and 5 present typical field SEIs
201 for submicron and supermicron aerosol particles collected at the ATTO and Manaus sites,
202 respectively, where the chemical species comprising each particle is indicated. Ammonium sulfate
203 and SOA particles are dominant in the sub- and super-micron aerosol fractions collected at the
204 ATTO site with some mineral particles and aged sea-salts in the supermicron fractions, whereas
205 the aerosol samples collected at the Manaus site are composed of various types of particles of
206 anthropogenic and/or natural origin. Figs. S1-S7 of Supporting Information present typical SEIs
207 for all the samples with identified chemical species on the SEIs, which helps briefly illustrate the
208 different features of the samples collected at the ATTO and Manaus sites. The characteristics of
209 the particle types observed in the ATTO and Manaus samples are described in the following.

210

211 **3.1.1. Secondary organic aerosol (SOA) particles**

212 In this study, SOA particles were observed frequently in both the ATTO and Manaus
213 samples, even though pure SOA particles were rare and most of them were mixed internally with

214 other species, such as ammonium sulfate, K-rich salt, reacted sea-salts, etc. The SOA particles over
215 the Amazon rainforest are formed by the oxidation of biogenic volatile organic compounds
216 (Jimenez et al., 2009; Hallquist et al., 2009; Martin et al., 2010a; Andreae et al., 2018) and are the
217 major constituents of particulate matter (PM), particularly for submicron ambient PM (Pöschl, et
218 al., 2010; Martin et al., 2010b; Chen et al., 2015). In the SEI images, pure SOA droplet particles
219 appear gray in contrast and have a circular shape, as shown in Fig. 6(a). As TEM grid films (with
220 90% C content) and SOAs are composed mainly of carbon and oxygen, the SOA aerosols appear
221 gray on the TEM grids because of their similar secondary and backscattered electron yields to
222 those of the TEM grid (Goldstein et al., 2003; Maskey et al., 2010). As TEM grids are hydrophobic
223 due to the thin carbon layer over the Formvar film, the aqueous droplet aerosols appear circular on
224 the TEM grids (Eom et al., 2014; Maskey et al., 2010), suggesting that SOAs were collected as
225 aqueous droplets at the time of particle sample collection. Recent studies also reported that most
226 submicron SOA particles in the Amazon basin are water soluble organic aerosols (WSOAs) rather
227 than semi-solid or solid aerosols under the background conditions that are typically met during the
228 wet season (Bateman et al., 2016, 2017). The X-ray spectrum of a typical pure SOA, as shown in
229 Fig. 6(a), showed considerably higher levels of the C X-ray peak intensity compared to that from
230 the Formvar/carbon film of the TEM grids, resulting in the unambiguous identification of SOA
231 particles based on their SEIs and X-ray spectral data.

232

233 **3.1.2. Ammonium sulfate (AS) particles**

234 Ammonium sulfate particles were observed abundantly in the ATTO and Manaus samples,
235 mostly as mixtures with secondary organics. Ammonium sulfate particles appear bright and
236 crystalline on the SEIs before the X-ray measurements, both for pure airborne and standard
237 ammonium sulfate particles, as shown in Figs. 6(b) and 6(f), respectively. The standard ammonium
238 sulfate particles were deposited on TEM grids by the nebulization of a 0.1 M ammonium sulfate
239 solution. As shown in the inset in Figs. 6(b) and 6(f), after the X-ray measurements, they show
240 somewhat darkened SEIs with black holes, due to electron beam damage, at the places where the
241 electron beam hits. As the ammonium sulfate particles are electron beam-sensitive (Geng et al.,
242 2010; Worobiec et al., 2003; Huang and Turpin, 1996), their X-ray spectral signature is the
243 presence of a significant S X-ray peak, as shown for both pure airborne and standard particles on
244 the TEM grids. The N X-ray peak was often not detected, particularly for small particles because

245 the NH_4^+ moiety is especially prone to damage by electron beams.

246 Ambient urban and rural sulfates act as a sink for ammonia, of which the sources are largely
247 animal waste, fertilizer application, soil release, and industrial emissions. The most common form
248 is ammonium sulfate. On the other hand, if ammonia is scarce in the air, sulfates would be in more
249 acidic forms, such as NH_4HSO_4 or H_2SO_4 (Millstein et al., 2008). The acidic $\text{NH}_4\text{HSO}_4/\text{H}_2\text{SO}_4$
250 particles have been reported to be more hygroscopic than pure ammonium sulfate (Pósfai et al.,
251 1998). Hence, they can be spread over the collecting substrate (Formvar/carbon film). In addition,
252 acidic sulfate particles can have unique halo rings in their morphology (Buseck and Pósfai, 1999).
253 The crystalline structure of the ammonium sulfate-containing particles observed in this study
254 suggests that they are sulfates fully neutralized with ammonia. In addition, the Raman spectra of
255 airborne particles exhibiting this morphology were obtained on a single particle basis to confirm
256 that they are ammonium sulfate. As shown in Fig. 7(a), the Raman peak at 975 cm^{-1} of the airborne
257 particle is characteristic of ammonium sulfate (Ling and Chan, 2007), which was also confirmed
258 by Raman spectroscopy on standard ammonium sulfate particles. Characteristic Raman peaks for
259 NH_4HSO_4 , K_2SO_4 , $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$, CaSO_4 , Na_2SO_4 , and $\text{MgSO}_4 \cdot x\text{H}_2\text{O}$ ($x = 1-7, 11$) were reported
260 to be at 1010 and 1042, 983, 1008, 1014 and 1025, 992, and 984-1046 cm^{-1} , respectively (Fung
261 and Tang, 1988; Wang et al., 2006; Mabrouk et al., 2013; Prieto-Taboada et al., 2014). The sloping
262 baseline in the Raman spectrum of the airborne particle was attributed to the fluorescence from
263 organics, indicating the presence of organic compounds (Sobanska et al., 2012), probably from
264 SOA. For aerosols collected on Ag foil at the ATTO site on June 10, 2014, ammonium sulfate is
265 the major species with some organics and minerals such as kaolinite, for the bulk aerosols in the
266 size range of 0.25 – 0.5 μm , based on their X-ray, attenuated total reflectance-FTIR (ATR-FTIR),
267 and Raman spectra (Fig. 7(b)). A study of the samples collected in the central Amazon Basin during
268 the wet season from February to March 2008 reported that ammonium was not sufficient to fully
269 neutralize sulfates so that ammonium bisulfate would be present in the Amazon rainforest (Chen
270 et al., 2015), whereas other studies reported that sulfates are sufficiently neutralized with ammonia
271 in the fine and coarse fractions during both the wet and dry seasons (Andreae et al., 2015; Martin
272 et al., 2010b; Fuzzi et al., 2007; Mace et al., 2003). The different results may be due to different
273 sampling places and seasons. In this study, ammonium sulfate is dominant over ammonium
274 bisulfate.

275 Previous studies have shown that the sulfate aerosols over the Amazon forest are

276 predominantly from marine and terrestrial biogenic sources, with comparable contributions from
277 marine and terrestrial biogenic emissions (Andreae et al., 1990). Sulfate originates from biogenic
278 sources in the rainforest, i.e., dimethyl sulfide (DMS), H₂S, and CS₂ emitted by plants and
279 microorganisms, which can be oxidized to sulfate (Andreae et al., 1990, 2015; Martin et al., 2010b).
280 The rainforest ecosystem in the central Amazon can act as a source of DMS to the atmosphere
281 throughout the year (Jardine et al., 2015). Several studies have reported that marine DMS
282 transported from the Atlantic Ocean contributes significantly to the sulfate levels in the Amazon
283 basin (Gregory et al., 1986; Andreae et al., 1990; Martin et al., 2010a). In addition, there is some
284 sulfate from long-range transport across the Atlantic including that from the African volcanic
285 emissions (Saturno et al., 2018) and minor upwind anthropogenic sources.

286 The nitrogen cycle is essential for organisms and some bacteria to fix the gaseous N₂ in the
287 air to NH₄⁺ for their own biosynthetic processes (Kim and Rees, 1994; Bazzaz, 1998; Kellerhals
288 et al., 2010). In addition, some microorganisms produce enzymes to release nitrogen as NH₄⁺
289 during the nitrogen mineralization process, which is important in tropical rain forest soils, where
290 dead plants and animal matter accumulate continuously (Wright, 1996; Neill et al., 1999). A high
291 level of NH₄⁺ in tropical rain contributes significantly to the nitrogen influx in the rainforest soils
292 (Jordan et al., 1982). The NH₄⁺ species can be evaporated as gaseous NH₃ from surface soils,
293 particularly from leaf litter, resulting in a strong NH₃ emission source as well as stomatal NH₃
294 emission of plants as another natural source in forest ecosystems (Sutton et al., 2009, 2013; Hansen
295 et al., 2017). On the other hand, NH₄⁺ species in rainforest soils might become airborne
296 immediately after rainfall, similar to the way that airborne organic particles are produced directly
297 from soils by raindrop impaction (Wang et al., 2016). The wetness of forest surfaces is significant
298 in controlling both the deposition and emission of atmospheric NH₃ (Hansen et al., 2015). As
299 ammonium sulfate-containing particles were also observed abundantly in the samples collected at
300 Manaus site, they were influenced strongly by the surrounding Amazonian forest and/or generated
301 by anthropogenic activities in the urban environment. In the urban environment, anthropogenic
302 ammonium sulfate is mainly formed by gaseous reactions among SO₂ emitted from coal-fired
303 plants and industrial activities, NH₃ emitted from human and animal activities and fertilization in
304 the fields, and oxidants (e.g., O₃ and OH radical) (Li et al., 2016; Geng et al., 2017).

305

306 **3.1.3. SOA and AS mixture particles**

307 In this study, most airborne submicron SOAs were observed to be internally mixed with
308 ammonium sulfate, particularly for the samples collected at the ATTO site. Figure 6(c) shows that
309 a typical SOA and AS mixture particle has the crystalline, bright ammonium sulfate moiety in the
310 center surrounded by circular, grey organic species. The circular morphology of the organic species
311 strongly suggests that the organic species are SOAs, as stated above. As efflorescence and
312 deliquescence relative humidity (ERH and DRH) of ammonium sulfate species are 30-40% and
313 80%, respectively (Yeung and Chan, 2010) and the ambient RH was always above 55% during
314 sampling for the ATTO and Manaus samples, the ammonium sulfate would be mostly in aqueous
315 droplets at the time of sample collection, rather than in crystalline form, as indicated by their
316 overall circular shape. When the particle samples were under dry conditions either during sample
317 storage or in the vacuum chamber of the SEM instrument, the ammonium sulfate species
318 crystallized, resulting in core-shell structures of organic and inorganic mixture aerosols. When the
319 ambient RH is low enough to make the ammonium sulfate species crystallize in the atmosphere,
320 the organic and inorganic mixture aerosols would similarly be present as core-shell structures.

321 Some of the SOA and AS mixture particles were also mixed with K-salts. As shown in Figs.
322 6(d) and 6(e), their morphology was similar to that of the SOA and AS mixture particles, but their
323 X-ray spectra revealed the presence of K and an elevated S level compared to those of the SOA
324 and AS mixture particles, suggesting that the K is associated mostly with SO_4^{2-} . The shoulder
325 Raman peak of the airborne ammonium sulfate particle at 982 cm^{-1} (Fig. 7(a)), which is indicative
326 of the K_2SO_4 moiety (Mabrouk et al., 2013), also suggests that the K-salts are most probably K_2SO_4 .
327 A previous study reported that small K-salt-rich particles can act as seeds for SOA formation in
328 the Amazon basin and K-salts are present ubiquitously in Amazonian SOAs with their content
329 being higher in the morning hours and for smaller SOAs (Pöhlker et al., 2012). On the other hand,
330 among the 843 submicron SOA and/or ammonium sulfate particles collected at the ATTO site,
331 only 31% contained K-salts, which is probably because the samples were collected in the afternoon
332 and/or the analyzed particles were larger than $0.25\text{ }\mu\text{m}$ so that the K-salt content may be below the
333 detection limit of EDX ($\sim 0.1\text{ wt. \%}$). In the Manaus samples, a total of 199 submicron SOA and/or
334 ammonium sulfate particles were observed, of which approximately 40% contained K-salts,
335 suggesting that the Manaus samples were influenced strongly by the surrounding rainforest as
336 supported by the backward trajectories (Figs. 2(e)-(g)), where K-salts may be mostly of a biogenic
337 origin in the rainforest. The organic moiety is often mixed internally with aged sea-salts, mineral

338 dust, and PBAs, which will be described below.

339

340 **3.1.4. Mineral dust particles**

341 The typical mineral dust particles include aluminosilicate, quartz (SiO_2), calcite (CaCO_3),
342 dolomite ($\text{CaMg}(\text{CO}_3)_2$), and TiO_2 (Geng et al., 2009, 2011). They appear irregular and bright on
343 the SEIs (Fig. 8). Various types of mineral particles from Saharan dust contribute significantly to
344 the nutrient cycles in the Amazon rainforest (Talbot et al., 1990; Abouchami et al., 2013; Rizzolo
345 et al., 2017). Mineral dust tends to provide reactive surfaces for heterogeneous reactions with trace
346 atmospheric gases, such as SO_2 and NO_x , leading to chemical modifications of the particles that
347 ultimately affect the atmospheric chemical balance and photochemical cycle (Sullivan et al., 2007;
348 Chen et al., 2011). Modification of the physicochemical properties of particles can alter their
349 optical, chemical, and hygroscopic properties (Sullivan et al., 2007; Geng et al., 2014). If some
350 components in them (particularly Ca-containing species) react with airborne SO_2 and NO_x in the
351 presence of moisture or with “secondary acids”, such as H_2SO_4 , HNO_3 , and HCl , they are regarded
352 as reacted or aged ones. The reacted/aged ones can contain either nitrates, sulfates, or both (Geng
353 et al., 2014, 2017; Li et al., 2016). In the Amazon samples, almost all the mineral dust particles
354 were aged ones, as shown in Fig. 8, where an aluminosilicate and a carbonate/silica mixture
355 particle containing sulfate are shown. The S-containing aged mineral dust particles outnumbered
356 the N- and both N- and S-containing ones for the ATTO samples, whereas they were comparable
357 to both the N- and S-containing ones for the Manaus samples, as shown in Fig. 9(a), indicating the
358 predominance of sulfates over nitrates for the reaction of mineral dust particles at the ATTO site
359 and somewhat significant influence of nitrates at the Manaus site. The nitrates for the Manaus
360 samples may be formed from nitrogen oxides emitted from the vehicles and coal-fired power plants
361 (Geng et al., 2014, 2017; Li et al., 2016). The mineral particles might be mixed or covered with
362 SOA and/or ammonium sulfate and gradually become aged under a high RH in the rainforest.
363 Some mineral particles were also mixed with sea salt particles during their transport to the
364 Amazonian area across the Atlantic Ocean.

365

366 **3.1.5. Reacted (aged) sea-salts**

367 The fresh sea-salt particles can react with gas-phase sulfur and nitrogen oxides to contain
368 sulfate and nitrate, respectively (ten Brink, 1998). During the process, chlorine may be removed

369 completely if the reaction is complete (Laskin et al., 2003). All sea-salt particles observed in the
370 ATTO and Manaus samples were reacted ones. Figure 10(a) shows the X-ray spectrum, atomic
371 concentration, and SEI of a typical aged sea salt particle, where the presence of Na, Mg, and Cl
372 indicates its marine origin (Geng et al., 2014) and the presence of S and C indicates that it is mixed
373 with sulfates and organics. The irregular and somewhat bright SEI is typical of the aged sea-salt
374 particles. As shown in Fig. 9(b), S-containing sea-salts outnumbered N- and both N- and S-
375 containing ones for the ATTO samples, indicating the predominance of sulfates over nitrates for
376 the reaction of sea-salt particles. The S-containing particles are comparable in abundance to both
377 the N- and S-containing ones at the Manaus site (Fig. 9(b)). The sea-salt particles may also become
378 mixed with ammonium sulfate over the rainforest and become S-containing ones. Among overall
379 275 reacted sea-salts containing sulfate/nitrate and organics, 71% of them were mixed with K-salts,
380 as shown in Fig. 10(b). The presence of K-salts in the reacted sea-salt particles indicates that the
381 mixing of the K-salts of biogenic origin would happen in the rainforest during long-range transport
382 because of the minimal biomass burning influence during the wet season, although it was reported
383 that a strong biomass burning smoke was transported from Africa to South America during the wet
384 season (Baars et al., 2011) so that a part of the observed K-salts could be associated with biomass
385 burning aerosols from Africa. In addition, several elongated CaSO_4 particles, as shown in Fig.
386 10(c), were detected in both ATTO and Manaus samples, all of which contain a small amount of
387 Na. Their elongated shape and the presence of Na strongly indicates that they were from the sea,
388 possibly the Atlantic Ocean, not from the soil (Eom et al., 2016).

389

390 **3.1.6. Primary biological aerosol (PBA) particles**

391 PBA particles like fungal spores can be identified easily based on their unique morphology
392 and the presence of their characteristic chemical elements (Geng et al., 2009; Martin et al., 2010a;
393 Pöschl et al., 2010). The PBA particle shown in Fig. 11(a) has a unique oval morphology and the
394 majority of C together with the characteristic small amounts of N, P, S, Cl, and K. PBA particles
395 are relatively large (diameter $> 2.0 \mu\text{m}$) so that they are abundant in stage 4 samples ($2.0 \mu\text{m} <$
396 $\text{diameter} < 4.0 \mu\text{m}$), particularly at the ATTO site. Figure 12 shows two image fields of stage 4
397 samples collected at the ATTO and Manaus sites, where the PBA particles have various types of
398 morphology and many of them are mixed with SOA. The abundant observation of PBA particles
399 in the stage 4 sample of the Manaus site suggests the transport of the PBA particles from the

400 rainforest to the urban area. Supermicron PBA particles were reported to be abundant over the
401 Amazon (China et al., 2016; Moran-Zuloaga et al., 2017; Gilardoni et al., 2011). PBA particles can
402 be pollen, bacteria, fungal and fern spores, viruses, and fragments of plants and animals emitted
403 directly from the rainforest, showing a range of morphologies, and comprise the largest fraction of
404 the coarse mass (Martin et al., 2010a). PBA particles appear to be the most efficient and abundant
405 ice nuclei (Pöschl et al., 2010; Tobo et al., 2013; Haga et al., 2014). In addition, the release of
406 nano- and submicron particles from fungal spores under high relative humidity can contribute to
407 new particle formation and potentially affect cloud formation in the Amazon Basin (China et al.,
408 2016).

409

410 **3.1.7. Carbon-rich particles from combustion sources**

411 Carbon-rich particles, such as soot, tarballs, and char or coal dust, which contain more than
412 90% C and O with the C content being dominant over O in low-Z particle EPMA analysis (Geng
413 et al., 2009, 2010, 2014), were observed frequently in the Manaus samples, whereas they were rare
414 in the ATTO samples. Based on the characteristic morphology of carbon-rich particles, soot
415 aggregates of fractal-like chain structures (Fig. 11(b)), tarballs of separate spherules (Fig. 11(c)),
416 and chars of irregular-shaped carbon (Fig. S5, SM1-2) could be differentiated straightforwardly
417 from each other (Geng et al., 2010, 2014; Li et al., 2016). The soot aggregates formed via a
418 vaporization-condensation mechanism during the combustion processes vary in size from sub to
419 several micrometers (Chen et al., 2005, 2006). Once airborne, the complex microstructure of the
420 soot aggregates may provide active sites for the deposition of organics and other chemical species,
421 such as sulfates (Pósfai et al., 1999; Zhang et al., 2008), as revealed by the presence of S in Fig.
422 11(b). This results in aged soot aggregates that become compact with considerable restructuring
423 and shrinkage (Zhang et al., 2008). Tarballs, which are a type of brown carbon (Andreae and
424 Gelencsér, 2006; Laskin et al., 2015), usually have high C, N, and O contents with a spherical
425 morphology (Fig. 11(c)), strongly indicating their formation during biomass combustion processes
426 (Pósfai et al., 2003, 2004). Char appears compact and irregular in the SEIs; it is often mixed with
427 minor inorganic species, such as K and S, and is regarded as the incomplete combustion remnants
428 of liquid or solid carbonaceous fuel materials that have undergone carbonization during
429 combustion (Chen et al., 2006). Only one soot particle was observed in all the ATTO samples,
430 whereas soot, tarball, and char or coal dust particles were abundant in the submicron Manaus

431 samples, suggesting that the ATTO samples are barely affected by the anthropogenic carbon-rich
432 particles generated in Manaus.

433

434 **3.1.8. Fly ash particles**

435 As shown in Fig. 11(d), fly ash particles are glassy spheres, composed mainly of O, Si, and
436 Al with minor components, such as Fe and Ca, which can be identified clearly by their spherical
437 shape and bright contrast on SEIs. The fly ash particles are different from tarballs having only C
438 and O signals in their X-ray spectra, even though both are generated during the combustion
439 processes (Geng et al., 2017). They were observed only in the Manaus samples, reflecting their
440 anthropogenic origin.

441

442 **3.1.9. Heavy metal-containing particles**

443 Heavy metal-containing (HM) particles, such as Ni-, Ti-, Zn-, and Fe-containing ones,
444 appear bright and irregular on SEM images, as shown in Figs. 11(e) and (f), and were observed
445 mostly in the fine fraction with more than a half of them being Fe-containing particles both in the
446 ATTO and Manaus samples. The Fe-containing particles in the ATTO samples were observed to
447 be associated with SOA (and ammonium sulfate), as shown in Fig 11(f), indicating its mixing with
448 the species of a biogenic emission origin. Sahara mineral dust has been reported to be essential for
449 the nutrient cycles in the Amazon rainforest because many types of minerals are carried and
450 transported into the rainforest, in which Fe is one of the important micronutrients in a Fe-limited
451 rainforest (Rizzolo et al., 2017). Among all the mineral dust particles observed in the samples,
452 approximately 40% of them contain Fe. The HM particles can also be of anthropogenic origin:
453 emitted from the streets or road surface as brake dust, road paint particles, diesel exhaust particles,
454 construction materials, and/or car catalyst materials (Qiao et al., 2016).

455

456 **3.2. Relative abundances of particle types observed in the ATTO and Manaus samples**

457 Figure 13 shows the relative abundance of the nine different particle types observed in the
458 ATTO and Manaus samples. In the stage 1 samples (0.25-0.5 μm size) of SA1-SA4 collected at
459 the ATTO site, almost all the particles were SOA and AS mixtures. In the stage 2 samples (0.5-1.0
460 μm size), SOA and AS mixture particles were dominant for the SA2 and SA3 samples. In the stage
461 2 samples of samples SA1 and SA4, SOA and AS mixture particles were most abundant, followed

462 by pure ammonium sulfates, aged mineral dust, and pure SOA particles. In the stage 2 samples of
463 SA1-SA4, the summed contents of SOA and ammonium sulfate were 73%, 99%, 85%, and 82%,
464 respectively, suggesting that SOA and AS are the predominant species in submicron aerosols
465 collected at the ATTO site. The observation of abundant submicron SOAs, which constitute a
466 significant fraction of fine aerosol mass during the wet season at the rainforest, has been reported
467 (Chen et al., 2015; Gilardoni et al., 2011). Although ammonium sulfates were reported to be
468 present in significant quantities in the Amazon basin (Andreae et al., 2015; Chen et al., 2015;
469 Fraund et al., 2017), this study emphasizes the observation of the predominant submicron
470 ammonium sulfates mixed with SOA.

471 In the stage 2 samples, aged mineral dust and sea-salts for the SA1 sample, reacted sea salts
472 for the SA3 sample, and aged mineral dust and sea-salts for the SA4 sample were significantly
473 observed, suggesting that the samples from outside the Amazon rainforest have different influences
474 because the mineral dust and sea-salt particles were all aged ones. The influences from outside
475 were observed more clearly for supermicron aerosols at the ATTO site. In the stage 3 samples (1.0-
476 2.0 μm size) of SA1-SA4, reacted sea-salt particles and aged mineral dust particles were
477 abundantly observed, although the summed relative abundances of SOA and ammonium sulfate
478 were 28%, 43%, 58%, and 50%, respectively, indicating that SOA and ammonium sulfate are
479 abundant even in supermicron ATTO aerosols. As all the mineral dust particles were aged, they are
480 not of local origin, and the observation of a high content of aged mineral dust particles in sample
481 SA1 highlights the importance of long-range transatlantic transport (see the fast-moving air-
482 masses from the Atlantic Ocean for the SA1 sample (Fig. 2(a) although a single backtrajectory
483 cannot confirm their African origin of the aged mineral dust particles). Many studies have
484 examined the influence of the Saharan dust particles over the Amazon rainforest region, starting
485 with the measurements made during the ABLE-2B campaign (Talbot et al., 1990; Swap et al.,
486 1992). Mineral dust is imported most frequently to the rainforest in March and April (Martin et al.,
487 2010a; Moran-Zuloaga et al., 2018), which increases the ground-based soil dust element levels
488 significantly (Artaxo et al., 2013). In this study, approximately 70% and 90% of the reacted sea-
489 salt/mineral dust particles in the ATTO and Manaus samples, respectively, were either mixed or
490 coated with organic matter in the Amazon basin and/or during the long-range transport.

491 PBA particles, which are from the rainforest, are sometimes observed in the stage 3 samples
492 for the SA1-SA4 samples. In the stage 4 sample (2.0-4.0 μm size) of the SA4 sample, the most

493 abundant particle type was PBA, followed in order by reacted sea-salt, SOA, aged mineral dust,
494 and ammonium sulfate, where both PBA, SOA, and ammonium sulfate particles of a local origin
495 and the reacted sea-salt and mineral dust particles from the outside are considerably present. In
496 summary, the aerosols collected at the ATTO site are mostly SOA and ammonium sulfate of a local
497 origin in the submicron fraction, although some of the submicron sulfate are of marine and distant
498 origin, whereas aerosols of both local and distant origins are significant in the supermicron fraction.

499 The aerosols collected at the Manaus site were diverse compared to those at the ATTO site.
500 As shown in Fig. 13, for the stage 1 and 2 samples, SOA and ammonium sulfate particles, including
501 their mixture, were the major components for the SM1 and SM2 samples, whereas they were
502 relatively less abundant for the SM3 sample. As SOA and ammonium sulfate particles can be from
503 the surrounding rainforest areas in addition to local anthropogenic sources, samples SM1-SM3
504 collected at the Manaus site appear to be influenced from the outside in the order of samples SM1
505 > SM2 > SM3. In addition, considerable amounts of submicron carbonaceous particles were
506 observed, such as soot, char, and tarballs, which are of a local origin. The most abundant submicron
507 aerosols for sample SM3 were carbonaceous ones, indicating that the local influence to the samples
508 is in the order of SM3 > SM2 > SM1. Samples SM1 and SM2 were collected during and just after
509 a national holiday, respectively, when all the institutions (private and public) were closed during
510 that day so that the traffics were quite low, similar to a weekend or vacation period. Sample SM3
511 was collected during a regular working day, so that sample SM3 is the only sample actually
512 exposed to the high traffic of light vehicles in the area.

513 In supermicron Manaus aerosols, PBA particles, aged mineral dust, and reacted sea-salts in
514 addition to SOA and carbonaceous particles are abundant. In the stage 3 aerosols of the SM1
515 sample, the most abundant particles were reacted sea-salts, followed by aged mineral dust, SOA,
516 ammonium sulfate, and PBA particles, which also indicates the strong influence on the SM1
517 sample from the outside. In stage 3 aerosols of the SM2 sample, the most abundant particles were
518 SOA, followed in order by aged mineral dust, reacted sea-salt, PBA, carbonaceous particles, and
519 ammonium sulfate, which also indicates the strong influence on the SM2 sample from the
520 surrounding rainforest areas. In the stages 3 and 4 aerosols of the SM3 sample, the most abundant
521 particles were aged mineral dust (36% and 52% for stages 3 and 4, respectively), followed by
522 carbonaceous particles, PBA, SOA, and reacted sea-salt. As the aged sea-salt contents were
523 relatively low (10% and 4% for stages 3 and 4, respectively), most of the aged mineral dusts appear

524 to be of a local origin. Fly ash and heavy metal-containing particles of an anthropogenic local
525 origin were considerable in the Manaus samples, i.e., they were observed in the range of 3-7%
526 relative abundances in the Manaus samples. In particular, soot particles of an anthropogenic origin
527 were observed ubiquitously in all the Manaus samples. Among the samples, the aerosols collected
528 at the Manaus site were different. The SM1 sample was influenced most strongly from the outside,
529 including the surrounding rainforest and transatlantic transport. The SM2 sample has some
530 influences by local sources as well as from the outside. The SM3 sample contains mainly aerosols
531 of an anthropogenic local origin in the submicron fraction and some influences from the outside
532 in the supermicron fraction. Figures 2(e)-(f) show that the backward trajectories at heights of 100-
533 m and 500-m are further from the outside in the order of SM1 > SM2 > SM3, which is consistent
534 with the characteristics of submicron and supermicron aerosols of the SM1-SM3 samples.

535

536 **Conclusions**

537 In this study, aerosol samples collected in the Amazonian rainforest and Manaus, Brazil
538 during the 2012 wet season were investigated on a single particle basis using low-Z particle EPMA.
539 The aerosol particles were classified into nine particle types based on their morphology on SEIs
540 together with the elemental concentrations of a total of 3,162 individual particles: (i) secondary
541 organic aerosols (SOA), (ii) ammonium sulfate (AS) particles, (iii) SOA and AS mixture particles,
542 (iv) aged mineral dust, (v) reacted sea-salts, (vi) primary biological aerosol (PBA) particles, (vii)
543 carbon-rich or elemental carbon (EC) particles such as soot, tarball, and char, (viii) fly ash particles,
544 and (ix) heavy metal (HM, such as Fe, Zn, Ni, and Ti)-containing particles. For submicron aerosols
545 collected at the ATTO site, the SOA and AS mixture particles were predominant (50-94% in
546 relative abundance) with the summed contents of SOA and ammonium sulfate being 73-100%. In
547 contrast, in the supermicron aerosols at the ATTO site, aged mineral dust and sea-salt (37-70%) as
548 well as SOA and ammonium sulfate (28-58%) were abundant. PBAs were observed abundantly in
549 the PM₂₋₄ fraction (46%), and EC and fly ash particles were absent in all the fractions. An analysis
550 of a bulk PM_{0.25-0.5} aerosol sample collected at the ATTO site using RMS and ATR-FTIR showed
551 that ammonium sulfate, organics, and minerals are the major chemical species, which is consistent
552 with the EPMA results.

553 In the submicron aerosols collected in Manaus, either SOA and ammonium sulfate (17-
554 80%) or EC particles (6-78%) were dominant, depending on the samples. The supermicron

555 aerosols collected in Manaus consisted mainly of aged mineral dust, reacted sea-salts, PBA, SOA,
556 ammonium sulfate, and EC particles. SOA, ammonium sulfate, and PBAs were mostly of a
557 biogenic origin and EC and HM-containing particles were of an anthropogenic origin. The aged
558 mineral dust and reacted sea-salt particles, including mineral dust mixed with sea-salts probably
559 during long-range transatlantic transport, were abundant in the supermicron fractions at both sites.
560 The submicron aerosol at the ATTO site was influenced mainly by the emission from the rainforest
561 and its supermicron aerosols showed additional contributions from long-range transport, including
562 the Atlantic Ocean and Sahara Desert, whereas the aerosols collected in Manaus showed different
563 local and outside contributions among the samples. Among all the aged mineral dust and reacted
564 sea-salt particles, sulfate-containing ones outnumbered those containing nitrates and both nitrate
565 and sulfate in the ATTO samples, whereas N and S containing particles were comparable to sulfate-
566 only ones in the Manaus samples, indicating the different sources and formation mechanisms of
567 secondary aerosols, i.e., the predominant presence of sulfate at the ATTO site from biogenic
568 emissions and elevated influences of nitrates from anthropogenic activities at the Manaus site.

569

570 **Author contributions**

571 LW, RHMG, and CR designed the experiment and LW, XL, and HKK carried out the
572 measurements and analyzed the data. CGGB, CIY, RAFDS, and CP organized and performed the
573 samplings. LW, HG, RHMG, CGGB, AFLG, and CR interpreted the observations and LW, HG,
574 RHMG, CGGB, MOA, and CR drafted the paper.

575

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587

588 **References**

589 Abouchami, W., Nathe, K., Kumar, A., Galer, S. J. G., Jochum, K. P., Williams, E., Horbe, A. M.
590 C., Rosa, J. W. C., Balsam, W., Adams, D., Mezger, K., and Andreae, M. O., Geochemical
591 and isotopic characterization of the Bodele Depression dust source and implications for
592 transatlantic dust transport to the Amazon Basin, *Earth Planet. Sci. Lett.*, 380, 112-123,
593 doi:10.1016/j.epsl.2013.08.028, 2013.

594 Andreae, M. O.: Aerosols before pollution, *Science*, 315, 50–51, doi:10.1126/science.1136529,
595 2007.

596 Andreae, M. O. and Gelencser, A.: Black carbon or brown carbon? The nature of light-absorbing
597 carbonaceous aerosols, *Atmos. Chem. Phys.*, 6, 3131–3148, doi:10.5194/acp-6-3131-2006,
598 2006.

599 Andreae, M. O., Berresheim, H., Bingemer, H., Jacob, D. J., Lewis, B. L., Li, S.-M., and Talbot,
600 R. W.: The atmospheric sulfur cycle over the Amazon Basin, 2. Wet Season, *J. Geophys. Res.*,
601 95, 16813–16824, doi:10.1029/JD095iD10p16813, 1990.

602 Andreae, M. O., Acevedo, O. C., Araujo, A., Artaxo, P., Barbosa, C. G. G., Barbosa, H. M. J., Brito,
603 J., Carbone, S., Chi, X., Cintra, B. B. L., da Silva, N. F., Dias, N. L., Dias-Junior, C. Q., Ditas,
604 F., Ditz, R., Godoi, A. F. L., Godoi, R. H. M., Heimann, M., Hoffmann, T., Kesselmeier, J.,
605 Konemann, T., Kruger, M. L., Lavric, J. V., Manzi, A. O., Lopes, A. P., Martins, D. L.,
606 Mikhailov, E. F., Moran-Zuloaga, D., Nelson, B. W., Nolscher, A. C., Santos Nogueira, D., T.
607 F. Piedade, M., Pohlker, C., Poschl, U., Quesada, C. A., Rizzo, L. V., Ro, C.-U., Ruckteschler,
608 N., Sa, L. D. A., de Oliveira Sa, M., Sales, C. B., dos Santos, R. M. N., Saturno, J., Schongart,
609 J., Sorgel, M., de Souza, C. M., de Souza, R. A. F., Su, H., Targhetta, N., Tota, J., Trebs, I.,
610 Trumbore, S., van Eijck, A., Walter, D., Wang, Z., Weber, B., Williams, J., Winderlich, J.,
611 Wittmann, F., Wolff, S., and Yanez-Serrano, A. M.: The Amazon Tall Tower Observatory
612 (ATTO): overview of pilot measurements on ecosystem ecology, meteorology, trace gases,
613 and aerosols, *Atmos. Chem. Phys.*, 15, 10723–10776, doi:10.5194/acp-15-10723-2015, 2015.

614 Andreae, M. O., Afchine, A., Albrecht, R., Holanda, B. A., Artaxo, P., Barbosa, H. M. J., Borrmann,
615 S., Cecchini, M. A., Costa, A., Dollner, M., Futterer, D., Jarvinen, E., Jurkat, T., Klimach, T.,
616 Konemann, T., Knote, C., Kramer, M., Krisna, T., Machado, L. A. T., Mertes, S., Minikin, A.,

617 Pöhlker, C., Pöhlker, M. L., Pöschl, U., Rosenfeld, D., Sauer, D., Schlager, H., Schnaiter, M.,
618 Schneider, J., Schulz, C., Spanu, A., Sperling, V. B., Voigt, C., Walser, A., Wang, J., Weinzierl,
619 B., Wendisch, M., and Ziereis, H.: Aerosol characteristics and particle production in the upper
620 troposphere over the Amazon Basin, *Atmos. Chem. Phys.*, 18, 921-961, doi:10.5194/acp-18-
621 921-2018, 2018.

622 Artaxo, P., Fernandes, E. T., Martins, J. V., Yamasoe, M. A., Hobbs, P. V., Maenhaut, W., Longo,
623 K. M., and Castanho, A.: Large-scale aerosol source apportionment in Amazonia, *J. Geophys.*
624 *Res.*, 103(D24), 31837–31847, doi:10.1029/98JD02346, 1998.

625 Artaxo, P., Rizzo, L. V., Brito, J. F., Barbosa, H. M., Arana, A., Sena, E. T., Cirino, G. G., Bastos,
626 W., Martin, S. T., Andreae, M., O.: Atmospheric aerosols in Amazonia and land use change:
627 from natural biogenic to biomass burning conditions, *Faraday Discuss.*, 165, 203-235, doi:
628 10.1039/C3FD00052D, 2013.

629 Baars, H., Ansmann, A., Althausen, D., Engelmann, R., Artaxo, P., Pauliquevis, T. and Souza, R.:
630 Further evidence for significant smoke transport from Africa to Amazonia, *Geophys. Res.*
631 *Let.*, 38, L20802, doi:10.1029/2011GL049200, 2011.

632 Bateman, A. P., Gong, Z., Liu, P., Sato, B., Cirino, G., Zhang, Y., Artaxo, P., Bertram, A. K., Manzi,
633 A. O., Rizzo, L. V., Souza, R. A. F., Zaveri, R. A., and Martin, S. T.: Sub-micrometre
634 particulate matter is primarily in liquid form over Amazon rainforest, *Nat. Geosci.*, 9, 34–37,
635 doi:10.1038/ngeo2599, 2016.

636 Bateman, A. P., Gong, Z., Harder, T. H., de Sá, S. S., Wang, B., Castillo, P., China, S., Liu, Y.,
637 O'Brien, R. E., Palm, B. B., Shiu, H.-W., Cirino, G. G., Thalman, R., Adachi, K., Alexander,
638 M. L., Artaxo, P., Bertram, A. K., Buseck, P. P., Gilles, M. K., Jimenez, J. L., Laskin, A.,
639 Manzi, A. O., Sedlacek, A., Souza, R. A. F., Wang, J., Zaveri, R., and Martin, S. T.:
640 Anthropogenic influences on the physical state of submicron particulate matter over a tropical
641 forest, *Atmos. Chem. Phys.*, 17, 1759–1773, doi:10.5194/acp-17-1759-2017, 2017.

642 Bazzaz, F. A.: Tropical Forests in a Future Climate: Changes in Biological Diversity and Impact
643 on the Global Carbon Cycle, *Climatic Change*, 39, 317-336, doi:10.1023/A:1005359605003,
644 1998.

645 Buseck, P. R. and Pósfai, M.: Airborne minerals and related aerosol particles: Effects on climate
646 and the environment, *Proc. Natl. Acad. Sci. U.S.A.* 96, 3372-3379,
647 doi:10.1073/pnas.96.7.3372, 1999.

648 Chen, Y., Shah, N., Huggins, F. E., and Huffman, G. P.: Characterization of ambient airborne
649 particles by energy-filtered transmission electron microscopy, *Aerosol Sci. Technol.*, 39, 509–
650 518, doi:10.1080/027868291001402, 2005.

651 Chen, Y., Shah, N., Huggins, F.E., and Huffman, G.P.: Microanalysis of ambient particles from
652 Lexington, KY, by electron microscopy, *Atmos. Environ.*, 40, 651–663,
653 doi:10.1016/j.atmosenv.2005.09.036, 2006.

654 Chen, H. H., Navea, J. G., Young, M. A., and Grassian, V. H.: Heterogeneous Photochemistry of
655 Trace Atmospheric Gases with Components of Mineral Dust Aerosol, *J. Phys. Chem. A*, 115,
656 490-499, doi:10.1021/jp110164j, 2011.

657 Chen, Q., Farmer, D. K., Rizzo, L. V., Pauliquevis, T., Kuwata, M., Karl, T. G., Guenther, A., Allan,
658 J. D., Coe, H., Andreae, M. O., Pöschl, U., Jimenez, J. L., Artaxo, P., and Martin, S. T.:
659 Submicron particle mass concentrations and sources in the Amazonian wet season (AMAZE-
660 08), *Atmos. Chem. Phys.*, 15, 3687–3701, doi:10.5194/acp-15-3687-2015, 2015.

661 China, S., Wang, B., Weis, J., Rizzo, L., Brito, J., Cirino, G. G., Kovarik, L., Artaxo, P., Gilles, M.
662 K., Laskin, A.: Rupturing of Biological Spores As a Source of Secondary Particles in
663 Amazonia, *Environ. Sci. Technol.*, 50, 12179-12186, doi: 10.1021/acs.est.6b02896, 2016.

664 de Sá, S. S., Palm, B. B., Campuzano-Jost, P., Day, D. A., Hu, W., Isaacman-VanWertz, G., Yee, L.
665 D., Brito, J., Carbone, S., Ribeiro, I. O., Cirino, G. G., Liu, Y. J., Thalman, R., Sedlacek, A.,
666 Funk, A., Schumacher, C., Shilling, J. E., Schneider, J., Artaxo, P., Goldstein, A. H., Souza, R.
667 A. F., Wang, J., McKinney, K. A., Barbosa, H., Alexander, M. L., Jimenez, J. L., and Martin,
668 S. T.: Urban influence on the concentration and composition of submicron particulate matter
669 in central Amazonia, *Atmos. Chem. Phys.*, 18, 12185-12206, doi:10.5194/acp-18-12185-2018,
670 2018.

671 Eom, H.-J., Gupta, D., Li, X., Jung, H.-J., Kim, H., and Ro, C.-U.: Influence of Collecting
672 Substrates on the Characterization of Hygroscopic Properties of Inorganic Aerosol Particles,
673 *Anal. Chem.*, 86, 2648–2656, doi:10.1021/ac4042075, 2014.

674 Eom, H.-J., Gupta, D., Cho, H.-R., Hwang, H. J., Hur, S. D., Gim, Y., and Ro, C.-U.: Single-
675 particle investigation of summertime and wintertime Antarctic sea spray aerosols using low-
676 Z particle EPMA, Raman microspectrometry, and ATR-FTIR imaging techniques, *Atmos.*
677 *Chem. Phys.*, 16, 13823-13836, 10.5194/acp-16-13823-2016, 2016.

678 Fan, J., Rosenfeld, D., Zhang, Y., Giangran, S. E., Li, Z., Machado, L. A. T., Martin, S. T., Yang,

679 Y., Wang, J., Artaxo, P., Barbosa, H. M. J., Braga, R. C., Comstock, J. M., Feng, Z., Gao, W.,
680 Gomes, H. B., Mei, F., Pöhlker, C., Pöhlker, M. L., Pöschl, U., and de Souza, R. A. F.:
681 Substantial convection and precipitation enhancements by ultrafine aerosol particles, *Science*,
682 359, 411-418, 2018.

683 Fraund, M., Pham, D. Q., Bonanno, D., Harder, T. H., Wang, B., Brito, J., Sá, S. S. de, Carbone,
684 S., China, S., Artaxo, P., Martin, S. T., Pöhlker, C., Andreae, M. O., Laskin, A., Gilles, M., K.,
685 and Moffet, R. C.: Elemental Mixing State of Aerosol Particles Collected in Central
686 Amazonia during GoAmazon2014/15, *Atmosphere*, 8, 173, doi:10.3390/atmos8090173, 2017.

687 Fung, K. H. and Tang, I. N.: Raman spectra of singly suspended supersaturated ammonium
688 bisulfate droplets, *Chem. Phys. Lett.*, 147, 509-513, doi:10.1016/0009-2614(88)85017-6,
689 1988.

690 Fuzzi, S., Decesari, S., Facchini, M. C., Cavalli, F., Emblico, L., Mircea, M., Andreae, M. O., Trebs,
691 I., Hoffer, A., Guyon, P., Artaxo, P., Rizzo, L. V., Lara, L. L., Pauliquevis, T., Maenhaut, W.,
692 Raes, N., Chi, X., Mayol-Bracero, O. L., Soto-García, L. L., Claeys, M., Kourttchev, I., Rissler,
693 J., Swietlicki, E., Tagliavini, E., Schkolnik, G., Falkovich, A. H., Rudich, Y., Fisch, G., and
694 Gatti, L. V.: Overview of the inorganic and organic composition of size-segregated aerosol in
695 Rondonia, Brazil, from the biomass-burning period to the onset of the wet season, *J. Geophys.*
696 *Res.* 112, D01201, doi:10.1029/2005JD006741, 2007.

697 Geng, H., Jung, H.-J., Park, Y., Hwang, H., Kim, H., Kim, Y.J., Sunwoo, Y., and Ro, C.-U.:
698 Morphological and chemical composition characteristics of summertime atmospheric
699 particles collected at Tokchok Island, Korea, *Atmos. Environ.* 43, 3364–3373,
700 doi:10.1016/j.atmosenv.2009.03.034, 2009.

701 Geng, H., Kang, S., Choel, M., Kim, H., and Ro, C.-U.: Characterization of individual
702 submicrometer aerosol particles collected in Incheon, Korea, by quantitative transmission
703 electron microscopy energy-dispersive X-ray spectrometry, *J. Geophys. Res.*, 115, D15306,
704 doi:10.1029/2009JD013486, 2010.

705 Geng, H., Ryu, J., Maskey, S., Jung, H.-J., and Ro, C.-U.: Characterisation of individual aerosol
706 particles collected during a haze episode in Incheon, Korea using the quantitative ED-EPMA
707 technique, *Atmos. Chem. Phys.* 11, 1327–1337, doi:10.5194/acp-11-1327-2011, 2011.

708 Geng, H., Hwang, H., Liu, X., Dong, S., and Ro, C.-U.: Investigation of aged aerosols in size
709 resolved Asian dust storm particles transported from Beijing, China, to Incheon, Korea, using

710 low-Z particle EPMA, *Atmos. Chem. Phys.*, 14, 3307–3323, doi:10.5194/acp-14-3307-2014,
711 2014.

712 Geng, H., Jin, C.-S., Zhang, D.-P., Wang, S.-R., Xu, X.-T., Wang, X.-R., Zhang, Y, Wu, L., and Ro,
713 C.-U.: Characterization of size-resolved urban haze particles collected in summer and winter
714 at Taiyuan City, China using quantitative electron probe X-ray microanalysis, *Atmos. Res.*,
715 190, 29-42, doi: 10.1016/j.atmosres.2017.02.005, 2017.

716 Gilardoni, S., Vignati, E., Marmer, E., Cavalli, F., Belis, C., Gianelle, V., Loureiro, A., and Artaxo,
717 P.: Sources of carbonaceous aerosol in the Amazon basin, *Atmos. Chem. Phys.*, 11, 2747-
718 2764, doi:10.5194/acp-11-2747-2011, 2011.

719 Goldstein, J., Newbury, D., Joy, D., Lyman, C., Echlin, P., Lifshin, E., Sawyer, L., and Michael, J.:
720 Scanning Electron Microscopy and X-ray Microanalysis, third ed., Kluwer Academic/Plenum
721 Publishers, New York, 2003.

722 Gregory, G. L., Harriss, R. C., Talbot, R. W., Rasmussen, R. A., Garstang, M., Andreae, M. O.,
723 Hinton, R. R., Browell, E. V., Beck, S. M., Sebacher, D. I., Khalil, M. A. K., Ferek, R. J., and
724 Harriss, S. V.: Air chemistry over the tropical forest of Guyana: *J. Geophys. Res.*, 91, 8603-
725 8612, 1986.

726 Haga, D. I., Burrows, S. M., Iannone, R., Wheeler, M. J., Mason, R. H., Chen, J., Polishchuk, E.
727 A., Pöschl, U., and Bertram, A. K.: Ice nucleation by fungal spores from the classes
728 *Agaricomycetes*, *Ustilaginomycetes*, and *Eurotiomycetes*, and the effect on the atmospheric
729 transport of these spores, *Atmos. Chem. Phys.*, 14, 8611-8630, doi:10.5194/acp-14-8611-
730 2014, 2014.

731 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J.,
732 Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T.,
733 Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W.,
734 McFiggans, G., Mentel, Th. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D.,
735 Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic
736 aerosol: current and emerging issues, *Atmos. Chem. Phys.*, 9, 5155–5236, doi:10.5194/acp-
737 9-5155-2009, 2009.

738 Hansen, K., Pryor, S. C., Boegh, E., Hornsby, K. E., Jensen, B., and Sørensen, L. L.: Background
739 concentrations and fluxes of atmospheric ammonia over a deciduous forest, *Agric. For.
740 Meteorol.*, 214-215, 380-392, doi:10.1016/j.agrformet.2015.09.004, 2015.

741 Hansen, K., Personne, E., Skjøth, C. A., Loubet, B., Ibrom, A., Jensen, R., Sørensen, L. L., and
742 Boegh, E.: Investigating sources of measured forest-atmosphere ammonia fluxes using two-
743 layer bi-directional modelling, *Agric. For. Meteorol.*, 237-238, 80-94,
744 doi:10.1016/j.agrformet.2017.02.008, 2017.

745 Huang, P. F. and Turpin, B.: Reduction of sampling and analytical errors for electron microscopic
746 analysis of atmospheric aerosols, *Atmos. Environ.*, 30, 4137-4148, 1996.

747 IBGE: Estimates of the resident population in Brazil and Federative Units, Brazilian Institute of
748 Geography and Statistics, 2017.

749 Jardine, K., Yañez-Serrano, A. M., Williams, J., Kunert, N., Jardine, A., Taylor, T., Abrell, L.,
750 Artaxo, P., Guenther, A., Hewitt, C. N., House, E., Florentino, A. P., Manzi, A., Higuchi, N.,
751 Kesselmeir, J., Behrendt, T., Veres, P. R., Derstroff, B., Fuentes, J. D., Martin, S. T., Andreae,
752 M. O.: Dimethyl sulfide in the Amazon rain forest, *Global Biogeochem. Cycles*, 29, 19–32,
753 doi:10.1002/2014GB004969, 2015.

754 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H.,
755 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M.,
756 Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin,
757 C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn,
758 M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J.
759 A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J.,
760 Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R.,
761 Takami, A., Miyoshi, T., Hatakeyama, S., Shimojo, A., Sun, J. Y., Zhang, Y. M., Dzepina, K.,
762 Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R.,
763 Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.:
764 Evolution of organic aerosols in the atmosphere, *Science*, 326, 1525–1529,
765 doi:10.1126/science.1180353, 2009.

766 Jordan, C., Caskey, W., Escalante, G., Herrera, R., Montagnini, F., Todd, R., and Uhl, C.: The
767 nitrogen cycle in a ‘Terra Firme’ rainforest on oxisol in the Amazon territory of Venezuela,
768 *Plant and Soil*, 67, 325-332, doi:10.1007/BF02182779, 1982.

769 Kellerhals, T., Brüttsch, S., Sigl, M., Knüsel, S., Gäggeler, H. W., and Schwikowski, M.:
770 Ammonium concentration in ice cores: A new proxy for regional temperature reconstruction?,
771 *J. Geophys. Res.*, 115, D16123, doi:10.1029/2009JD012603, 2010.

772 Kim, J. and Rees, D. C.: Nitrogenase and biological nitrogen fixation, *Biochemistry*, 33, 389–397,
773 1994.

774 Krejci, R., Ström, J., Reus, M. de, Sahle, W.: Single particle analysis of the accumulation mode
775 aerosol over the northeast Amazonian tropical rain forest, Surinam, South America, *Atmos.*
776 *Chem. Phys.*, 5, 3331–3344, 2005.

777 Laskin, A., Gaspar, D. J., Wang, W., Hunt, S. W., Cowin, J. P., Colson, S. D., and Finlayson- Pitts,
778 B. J.: Reactions at Interfaces as a Source of Sulfate Formation in Sea-Salt Particles, *Science*,
779 301, 340–344, doi:10.1126/science.1085374, 2003.

780 Laskin, A., Laskin, J., and Nizkorodov, S. A.: Chemistry of Atmospheric Brown Carbon, *Chem.*
781 *Rev.*, 115, 4335–4382, doi:10.1021/cr5006167, 2015.

782 Li, W., Shao, L., Zhang, D., Ro, C.-U., Hu, M., Bi, X., Geng, H., Matsuki, A., Niu, H., and Chen,
783 J.: A review of single aerosol particle studies in the atmosphere of East Asia: morphology,
784 mixing state, source, and heterogeneous reactions, *J. Clean. Prod.*, 112, 1330–1349,
785 doi:10.1016/j.jclepro.2015.04.050, 2016.

786 Lin, G., Sillman, S., Penner, J. E., and Ito, A.: Global modeling of SOA: the use of different
787 mechanisms for aqueous-phase formation, *Atmos. Chem. Phys.*, 14, 5451–5475,
788 doi:10.5194/acp-14-5451-2014, 2014.

789 Ling, T. Y., and Chan, C. K.: Formation and transformation of metastable double salts from the
790 crystallization of mixed ammonium nitrate and ammonium sulfate particles, *Environ. Sci. &*
791 *Techno*, 41, 8077–8083, 10.1021/es071419t, 2007.

792 Mabrouk, K. B., Kauffmann, T. H., Aroui, H., and Fontana, M. D.: Raman study of cation effect
793 on sulfate vibration modes in solid state and in aqueous solutions, *J. Raman Spectrosc*, 44,
794 1603–1608, doi: 10.1002/jrs.4374, 2013.

795 Mace, K. A., Artaxo, P., and Duce, R. A.: Water-soluble organic nitrogen in Amazon Basin aerosols
796 during the dry (biomass burning) and wet seasons, *J. Geophys. Res.*, 108(D16), 4512,
797 doi:10.1029/2003JD003557, 2003.

798 Martin, S. T., Andreae, M. O., Artaxo, P., Baumgardner, D., Chen, Q., Goldstein, A. H., Guenther,
799 A., Heald, C. L., Mayol-Bracero, O. L., McMurry, P. H., Pauliquevis, T., Pöschl, U., Prather,
800 K. A., Roberts, G. C., Saleska, S. R., Dias, M. A. S., Spracklen, D. V., Swietlicki, E., and
801 Trebs, I.: Sources and properties of Amazonian aerosol particles, *Rev. Geophys.*, 48, RG2002,
802 doi:10.1029/2008rg000280, 2010a.

803 Martin, S. T., Andreae, M. O., Althausen, D., Artaxo, P., Baars, H., Borrmann, S., Chen, Q., Farmer,
804 D. K., Guenther, A., Gunthe, S. S., Jimenez, J. L., Karl, T., Longo, K., Manzi, A., Müller, T.,
805 Pauliquevis, T., Petters, M. D., Prenni, A. J., Pöschl, U., Rizzo, L. V., Schneider, J., Smith, J.
806 N., Swietlicki, E., Tota, J., Wang, J., Wiedensohler, A., and Zorn, S. R.: An overview of the
807 Amazonian Aerosol Characterization Experiment 2008 (AMAZE-08), *Atmos. Chem. Phys.*,
808 10, 11415–11438, doi:10.5194/acp-10-11415-2010, 2010b.

809 Martin, S. T., Artaxo, P., Machado, L. A. T., Manzi, A. O., Souza, R. A. F., Schumacher, C., Wang,
810 J., Andreae, M. O., Barbosa, H. M. J., Fan, J., Fisch, G., Goldstein, A. H., Guenther, A.,
811 Jimenez, J. L., Pöschl, U., Silva Dias, M. A., Smith, J. N., and Wendisch, M.: Introduction:
812 Observations and Modeling of the Green Ocean Amazon (GoAmazon2014/5), *Atmos. Chem.*
813 *Phys.*, 16, 4785-4797, doi:10.5194/acp-16-4785-2016, 2016.

814 Maskey, S., Choël, M., Kang, S., Hwang, H., Kim, H., and Ro, C.-U.: The influence of collecting
815 substrates on the single-particle characterization of real atmospheric aerosols, *Anal. Chim.*
816 *Acta*, 658, 120-127, doi:10.1016/j.aca.2009.11.006, 2010.

817 Medeiros, A. S. S., Calderaro, G., Guimarães, P. C., Magalhaes, M. R., Morais, M. V. B., Rafee, S.
818 A. A., Ribeiro, I. O., Andreoli, R. V., Martins, J. A., Martins, L. D., Martin, S. T., and Souza,
819 R. A. F.: Power plant fuel switching and air quality in a tropical, forested environment, *Atmos.*
820 *Chem. Phys.*, 17, 8987-8998, doi:10.5194/acp-17-8987-2017, 2017.

821 Millstein, D. E., Harley, R. A., and Hering, S. V.: Weekly cycles in fine particulate nitrate, *Atmos.*
822 *Environ.* 42, 632–641, doi:10.1016/j.atmosenv.2007.10.010, 2008.

823 Möhler, O., Benz, S., Saathoff, H., Schnaiter, M., Wagner, R., Schneider, J., Walter, S., Ebert, V.,
824 Wagner, S.: The effect of organic coating on the heterogeneous ice nucleation efficiency of
825 mineral dust aerosols, *Environ. Res. Lett.*, 3, 025007, doi:10.1088/1748-9326/3/2/025007,
826 2008.

827 Moran-Zuloaga, D., Ditas, F., Walter, D., Saturno, J., Brito, J., Carbone, S., Chi, X., de Angelis, I.
828 H., Baars, H., Godoi, R. H. M., Heese, B., Holanda, B. A., Lavrič, J.V., Martin, S. T., Ming,
829 J., Pöhlker, M. L., Ruckteschler, N., Su, H., Wang, Y., Wang, Q., Wang, Z., Weber, B., Wolff,
830 S., Artaxo, P., Pöschl, U., Andreae, M. O., and Pöhlker, C.: Long-term study on coarse mode
831 aerosols in the Amazon rain forest with the frequent intrusion of Saharan dust plumes, *Atmos.*
832 *Chem. Phys.*, 18, 10055-10088, doi:10.5194/acp-18-10055-2018, 2018.

833 Neill, C., Piccolo, M. C., Melillo, J. M., Steudler, P. A., Cerri, C. C.: Nitrogen dynamics in Amazon

834 forest and pasture soils measured by ¹⁵N pool dilution, *Soil Biol. Biochem.*, 31, 567-572,
835 1999.

836 Palm, B. B., de Sá, S. S., Day, D.A., Campuzano-Jost, P., Hu, W., Seco, R., Sjostedt, S. J., Park,
837 J.-H., Guenther, A. B., Kim, S., Brito, J., Wurm, F., Artaxo, P., Thalman, R., Wang, J., Yee, L.
838 D., Wernis, R., Isaacman-VanWertz, G., Goldstein, A. H., Liu, Y., Springston, S. R., Souza,
839 R., Newburn, M. K., Alexander, M. L., Martin, S.T., and Jimenez, J. L.: Secondary organic
840 aerosol formation from ambient air in an oxidation flow reactor in central Amazonia, *Atmos.*
841 *Chem. Phys.*, 18, 467–493, doi:10.5194/acp-18-467-2018, 2018.

842 Pöhlker, C., Wiedemann, K. T., Sinha, B., Shiraiwa, M., Gunthe, S. S., Smith, M., Su, H., Artaxo,
843 P., Chen, Q., Cheng, Y. F., Elbert, W., Gilles, M. K., Kilcoyne, A. L. D., Moffet, R. C.,
844 Weigand, M., Martin, S. T., Poeschl, U., and Andreae, M. O.: Biogenic potassium salt
845 particles as seeds for secondary organic aerosol in the Amazon, *Science*, 337, 1075–1078,
846 doi:10.1126/science.1223264, 2012.

847 Pöhlker, M. L., Ditas, F., Saturno, J., Klimach, T., Hrabě de Angelis, I., Araùjo, A. C., Brito, J.,
848 Carbone, S., Cheng, Y., Chi, X., Ditz, R., Gunthe, S. S., Holanda, B. A., Kandler, K.,
849 Kesselmeier, J., Könemann, T., Krüger, O. O., Lavrič, J. V., Martin, S. T., Mikhailov, E.,
850 Moran-Zuloaga, D., Rizzo, L. V., Rose, D., Su, H., Thalman, R., Walter, D., Wang, J., Wolff,
851 S., Barbosa, H. M. J., Artaxo, P., Andreae, M. O., Pöschl, U., and Pöhlker, C.: Long-term
852 observations of cloud condensation nuclei over the Amazon rain forest – Part 2: Variability
853 and characteristics of biomass burning, long-range transport, and pristine rain forest aerosols,
854 *Atmos. Chem. Phys.*, 18, 10289-10331, doi:10.5194/acp-18-10289-2018, 2018.

855 Pöschl, U., Martin, S. T., Sinha, B., Chen, Q., Gunthe, S. S., Huffman, J. A., Borrmann, S., Farmer,
856 D. K., Garland, R. M., Helas, G., Jimenez, J. L., King, S. M., Manzi, A., Mikhailov, E.,
857 Pauliquevis, T., Petters, M. D., Prenni, A. J., Roldin, P., Rose, D., Schneider, J., Su, H., Zorn,
858 S. R., Artaxo, P., and Andreae, M. O.: Rainforest Aerosols as Biogenic Nuclei of Clouds and
859 Precipitation in the Amazon, *Science* 329, 1513-1516, doi:10.1126/science.1191056, 2010.

860 Pósfai, M., Xu, H., Anderson, and Buseck, P. R.: Wet and dry sizes of atmospheric aerosol particles:
861 An AFM-TEM study, *Geophys. Res. Lett.* 25, 1907–1910, doi:10.1029/98GL01416, 1998.

862 Pósfai, M., Anderson, J. R., Buseck, P. R., and Sievering, H.: Soot and sulfate aerosol particles in
863 the remote marine troposphere, *J. Geophys. Res.* 104, 21, 685–21, 693,
864 doi:10.1029/1999JD900208, 1999.

865 Pósfai, M., Simonics, R., Li, J., Hobbs, P. V., and Buseck, P. R.: Individual aerosol particles from
866 biomass burning in southern Africa: 1. Compositions and size distributions of carbonaceous
867 particles, *J. Geophys. Res.*, 108(D13), 8483, doi:10.1029/2002JD002291, 2003.

868 Pósfai, M., Gelencser, A., Simonics, R., Arato, K., Li, J., Hobbs, P. V., and Buseck, P. R.:
869 Atmospheric tarballs: Particles from biomass and biofuel burning, *J. Geophys. Res.*, 109,
870 D06213, doi:10.1029/2003JD004169, 2004.

871 Prieto-Taboada, N., Gomez-Laserna, O., Martinez-Arkarazo, I., Olazabal, M. A., and Madariaga,
872 J. M.: Raman Spectra of the Different Phases in the CaSO₄-H₂O System, *Anal. Chem.*, 86,
873 20, 10131-10137, doi: 10.1021/ac501932f, 2014.

874 Qiao, T., Zhao, M., Xiu, G., and Yu, J.: Simultaneous monitoring and compositions analysis of PM₁
875 and PM_{2.5} in Shanghai: implications for characterization of haze pollution and source
876 apportionment, *Sci. Total Environ.*, 557-558, 386-394, doi:10.1016/j.scitotenv.2016.03.095,
877 2016.

878 Rizzolo, J. A., Barbosa, C. G. G., Borillo, G. C., Godoi, A. F. L., Souza, R. A. F., Andreoli, R. V.,
879 Manzi, A. O., Sá, M. O., Alves, E. G., Pöhlker, C., Angelis, I. H., Ditas, F., Saturno, J., Moran-
880 Zuloaga, D., Rizzo, L. V., Rosário, N. E., Pauliquevis, T., Santos, R. M. N., Yamamoto, C. I.,
881 Andreae, M. O., Artaxo, P., Taylor, P. E., and Godoi, R. H. M.: Soluble iron nutrients in
882 Saharan dust over the central Amazon rainforest, *Atmos. Chem. Phys.*, 17, 2673-2687,
883 doi:10.5194/acp-17-2673-2017, 2017.

884 Saturno, J., Ditas, F., Penning de Vries, M., Holanda, B. A., Pöhlker, M. L., Carbone, S., Walter,
885 D., Bobrowski, N., Brito, J., Chi, X., Gutmann, A., Hrabe de Angelis, I., Machado, L. A. T.,
886 Moran-Zuloaga, D., Rüdiger, J., Schneider, J., Schulz, C., Wang, Q., Wendisch, M., Artaxo,
887 P., Wagner, T., Pöschl, U., Andreae, M. O., and Pöhlker, C.: African volcanic emissions
888 influencing atmospheric aerosols over the Amazon rain forest, *Atmos. Chem. Phys.*, 18,
889 10391-10405, doi:10.5194/acp-18-10391-2018, 2018.

890 Sobanska, S., Hwang, H., Choël, M., Jung, H., Eom, H., Kim, H., Barbillat, J., and Ro, C.-U.:
891 Investigation of the chemical mixing state of individual Asian Dust particles by the combined
892 use of Electron Probe X-ray Microanalysis and Raman Microspectrometry, *Anal. Chem.*, 84,
893 3145-3154, doi:10.1021/ac2029584, 2012.

894 Su, L., Yuan, Z., Fung, J. C. H., and Lau, A. K. H.: A comparison of HYSPLIT backward
895 trajectories generated from two GDAS datasets, *Sci. Total Environ.*, 506-507, 527-537,

896 doi:10.1016/j.scitotenv.2014.11.072, 2015.

897 Sullivan, R. C., Guazzotti, S. A., Sodeman, D. A., and Prather, K. A.: Direct observations of the
898 atmospheric processing of Asian mineral dust, *Atmos. Chem. Phys.*, 7, 1213-1236,
899 doi:10.5194/acp-7-1213-2007, 2007.

900 Sutton, M.A., Nemitz, E., Milford, C., Campbell, C., Erisman, J. W., Hensen, A., Cellier, P., David,
901 M., Loubet, B., Personne, E., Schjoerring, J. K., Mattsson, M., Dorsey, J. R., Gallagher, L.,
902 Horvath, M. W., Weidinger, T., Meszaros, R., Dämmgen, U., Neftel, A., Herrmann, B.,
903 Lehman, B. E., Flechard, C., and Burkhardt, J.: Dynamics of ammonia exchange with cut
904 grassland: synthesis of results and conclusions of the GRAMINAE integrated experiment.
905 *Biogeosciences* 6, 2907–2934. doi:10.5194/bg-6-2907-2009, 2009.

906 Sutton, M.A., Reis, S., Riddick, S.N., Dragosits, U., Nemitz, E., Theobald, M. R., Tang, Y. S.,
907 Braban, C. F., Vieno, M., Dore, A. J., Mitchell, R. F., Wanless, S., Daunt, F., Fowler, D.,
908 Blackall, T. D., Milford, C., Flechard, C. R., Loubet, B., Massad, R., Cellier, P., Personne, E.,
909 Coheur, P. F., Clarisse, L., Van Damme, M., Ngadi, Y., Clerbaux, C., Skjøth, C. A., Geels, C.,
910 Hertel, O., Wichink Kruit, R. J., Pinder, R. W., Bash, J. O., Walker, J. T., Simpson, D., Horvath,
911 L., Misselbrook, T. H., Bleeker, A., Dentener, F., and de Vries, W.: Towards a climate-
912 dependent paradigm of ammonia emission and deposition, *Philos. Trans. R. Soc. Lond. Ser.*
913 *B*, 368, doi:10.1098/rstb.2013.0166, 2013.

914 Swap, R., Garstang, M., Greco, S., Talbot, R., and Kållberg, P.: Saharan dust in the Amazon Basin,
915 *Tellus*, 44B, 133-149, 1992.

916 Talbot, R. W., Andreae, M. O., Berresheim, H., Artaxo, P., Garstang, M., Harriss, R. C., Beecher,
917 K. M., and Li, S. M.: Aerosol chemistry during the wet season in Central Amazonia: The
918 influence of long-range transport, *J. Geophys. Res.*, 95, 16,955-16,969, 1990.

919 ten Brink, H. M.: Reactive uptake of HNO₃ and H₂SO₄ in sea-salt (NaCl) particles, *J. Aerosol Sci.*,
920 29, 57-64, 10.1016/s0021-8502(97)00460-6, 1998.

921 Tobo, Y., Zhang, D., Matsuki, A., and Iwasaka, Y.: Asian dust parti- cles converted into aqueous
922 droplets under remote marine atmo- spheric conditions, *Proc. Natl. Acad. Sci. USA*, 96,
923 3396–3403, 2010.

924 Tobo, Y., Prenni, A. J., DeMott, P. J., Huffman, J. A., McCluskey, C. S., Tian, G., Pöhlker, C.,
925 Pöschl, U., and Kreidenweis, S. M.: Biological aerosol particles as a key determinant of ice
926 nuclei populations in a forest ecosystem, *J. Geophys. Res.*, 118, 10100-10110,

927 doi:10.1002/jgrd.50801, 2013.

928 Vekemans, B., Janssens, K., Vincze, L., Adams, F., and Van Espen, P.: Analysis of X-ray spectra
929 by iterative least squares (AXIL): New developments, *X-Ray Spectrom.* 23, 278-285, 1994.

930 Wang, A., Freeman, J. J., Jolliff, B. L., and Chou I-M.: Sulfates on Mars: A systematic Raman
931 spectroscopic study of hydration states of magnesium sulfates, *Geochim. Cosmochim.*
932 *Acta*, 70, 611, doi:10.1016/j.gca.2006.05.022, 2006.

933 Wang, B., Harder, T. H., Kelly, S. T., Piens, D. S., China, S., Kovarik, L., Keiluweit, M., Arey,
934 B. W., Gilles, M. K., and Laskin, A.: Airborne soil organic particles generated by precipitation,
935 *Nature Geosci.* 9, 433–437, doi:10.1038/ngeo2705, 2016.

936 Worobiec, A., de Hoog, J., Osan, J., Szaloki, I., Ro, C.-U., and Van Grieken, R.: Thermal
937 stability of beam sensitive atmospheric aerosol particles in electron probe microanalysis at
938 liquid nitrogen temperature, *Spectrochim. Acta B* 58, 479–496, 2003.

939 Wright, S. J.: Phenological responses to seasonality in tropical forest plants, 440-460, in *Tropical*
940 *Forest Plant Ecophysiology*, Mulkey S. S., Chazdon, R. L., and Smith A. P. (Eds), Springer,
941 1996.

942 Yeung, M. C., and Chan, C. K.: Water content and phase transitions in particles of inorganic and
943 organic species and their mixtures using Micro-Raman Spectroscopy, *Aerosol Sci. Technol.*,
944 4, 269-280, doi:10.1080/02786820903583786, 2010.

945 Zhang, R., Khalizov, A. F., Khalizov, J., Zhang, D., Xue, H., and McMurry, P. H.: Variability in
946 morphology, hygroscopicity, and optical properties of soot aerosols during atmospheric
947 processing, *Proc. Natl. Acad. Sci. U. S. A.*, 105, 10,291–10,296,
948 doi:10.1073/pnas.0804860105, 2008.

Figure 1. Location of sampling sites at the Brazilian Amazon basin: an urban site in Manaus (S 3°05.753' W 59°59.419') and a rainforest site at ATTO (S 02°647' W 58°59.992'). Map of South America (top left) with the region marked with a red rectangle and a map of the Amazonas state, Brazil (bottom left) also with the region of interest marked in red.

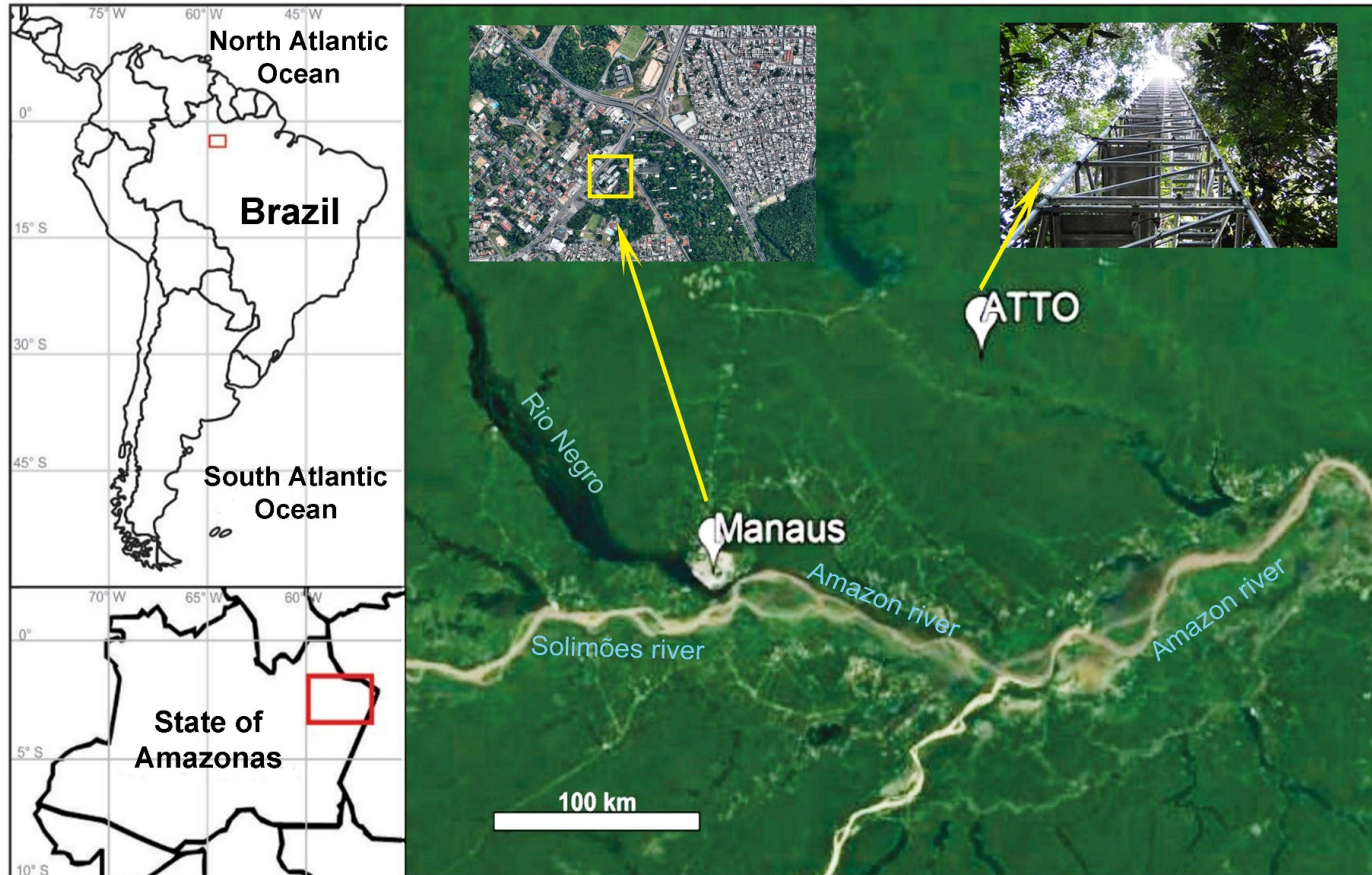
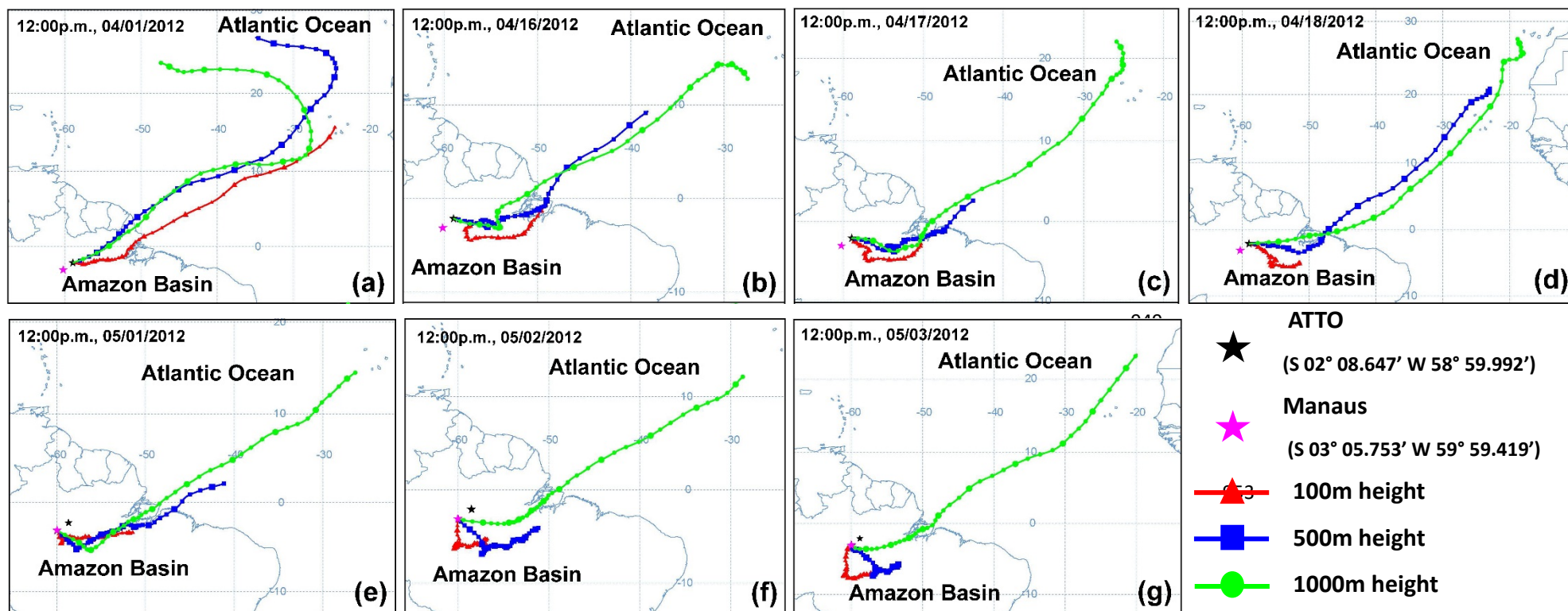


Figure 2. Ten-day (240 h) backward air mass trajectories at 100 m-, 500 m-, and 1000 m-receptor heights; (a)-(d) for the SA1-SA4 samples collected on April 1 and 16-18, 2012 at the ATTO site and (e)-(g) for the SM1-SM3 samples collected on May 1-3, 2012 at the Manaus site. HYbrid Lagrangian Single-Particle Integrated Trajectory (HYSPLIT) model available at the NOAA Air Resources Laboratory's web server (<http://www.arl.noaa.gov/ready/hysplit4.html>) was used



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Figure 3. Home-made sample holder for TEM grid samples in SEM/EDX measurements and a typical X-ray spectrum of the TEM grids.

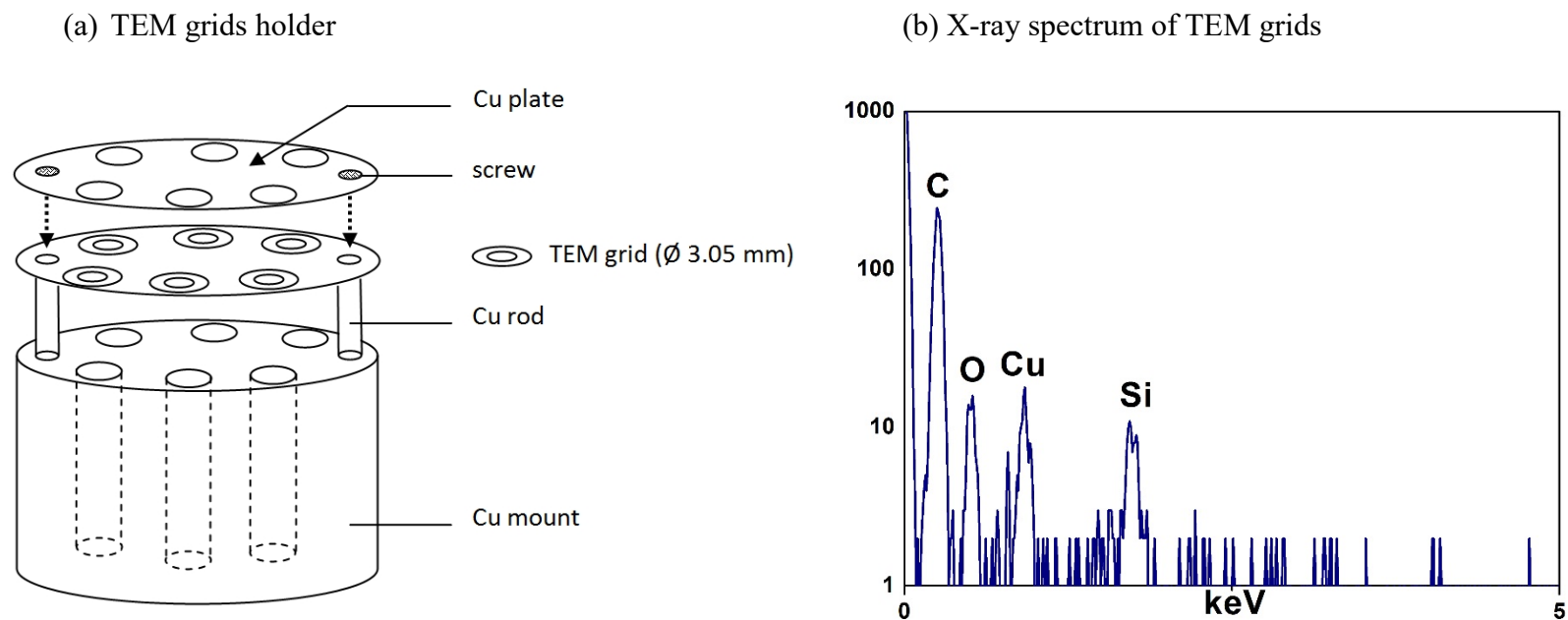


Figure 4. Typical SEM images of aerosol particles for (a) stage 1 ($PM_{0.25-0.5}$) of the SA2 sample, (b) stage 2 ($PM_{0.5-1.0}$) of the SA4 sample, (c) stage 3 (PM_{1-2}) of the SA1 sample, and (d) stage 4 ($PM_{2.0-4.0}$) of the SA4 sample, collected at the ATTO site. For convenience, ammonium sulfate, secondary organic aerosol, aluminosilicates, and reacted sea-salt are denoted as “AS”, “SOA”, “AlSi” and “rss”, respectively.

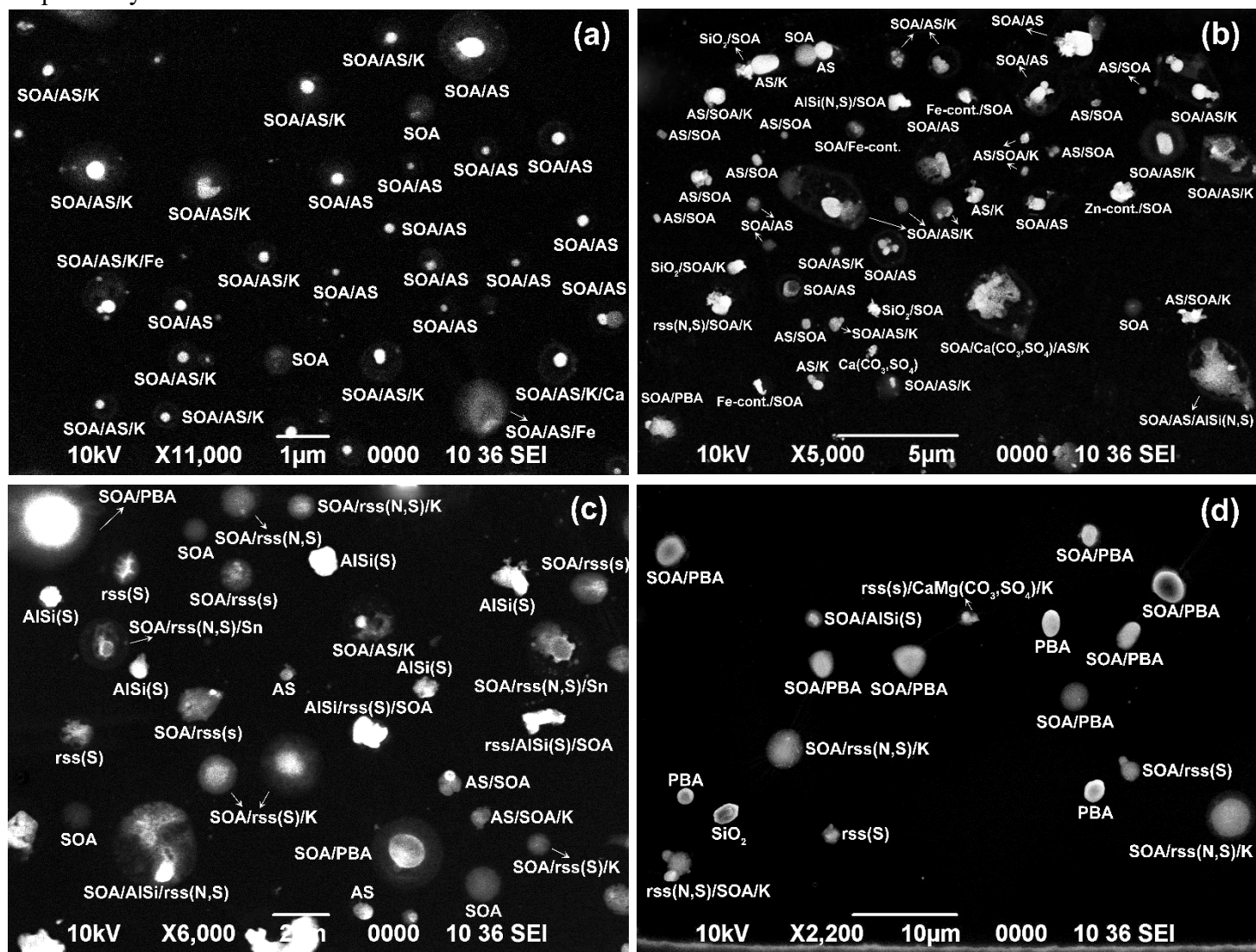


Figure 6. SEIs, X-ray spectra, and element atomic concentrations of SOA, ammonium sulfate (AS), and mixture particles. The inset images in (b), (c), and (f) show the beam damage on the particles after X-ray measurements.

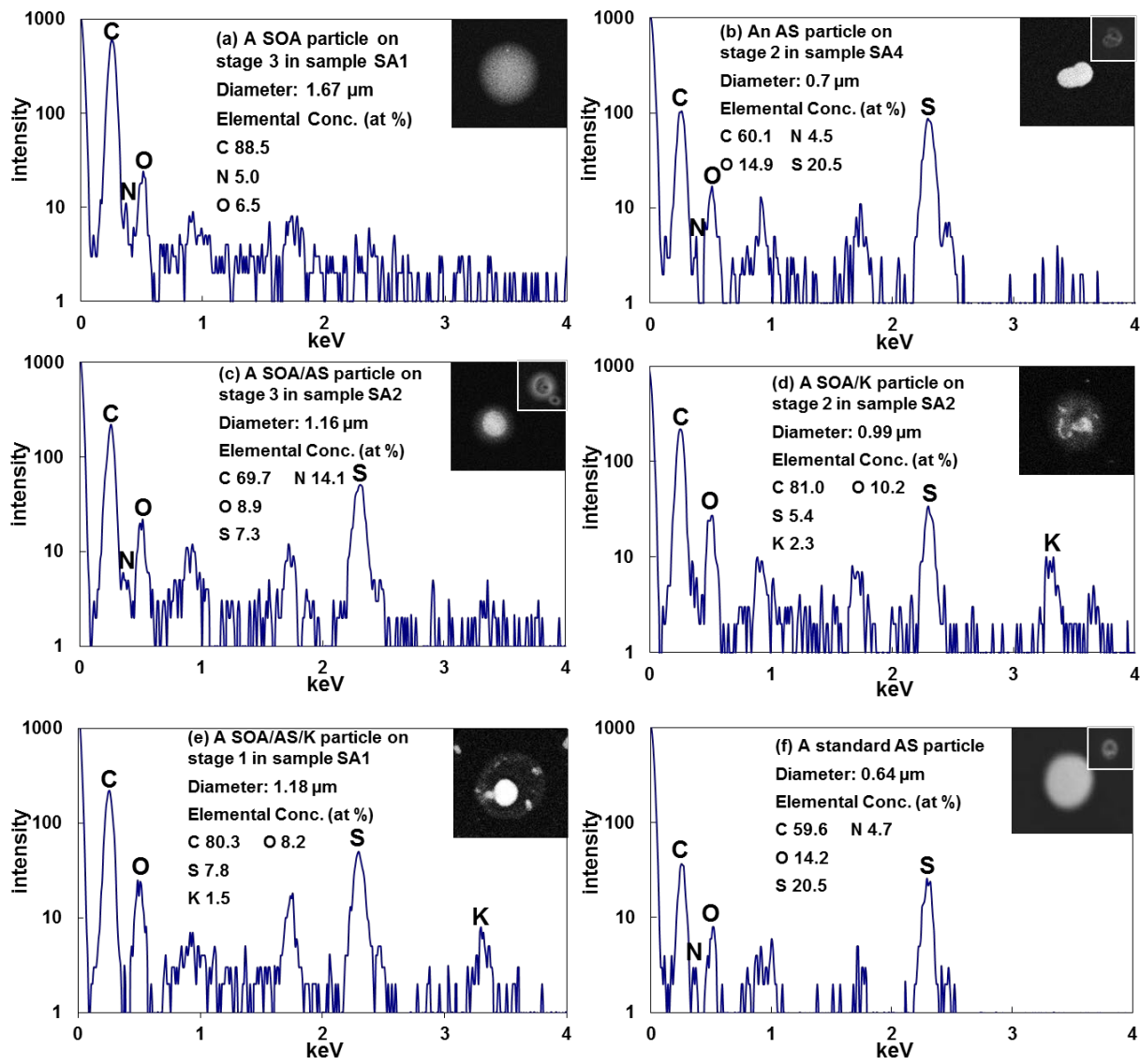


Figure 7. (a) Raman spectra of standard and airborne ammonium sulfate (AS) particles, which were rescaled for clarity. The inset SEI images are for standard and airborne AS particles where the scale bar is 1 μm . The shoulder peak of SO_4^{2-} at 982 cm^{-1} in the airborne AS particles is from K_2SO_4 ; (b) SEI, optical images, X-ray, ATR-FTIR, and Raman spectra of an overloaded $\text{PM}_{0.25-0.5}$ sample collected at ATTO site on June 10, 2014. X-ray, ATR-FTIR, and Raman spectra indicate that AS, organics, and minerals are the major components of the submicron sample. The sloping baseline in the airborne Raman spectrum is due to the fluorescence from organic compounds.

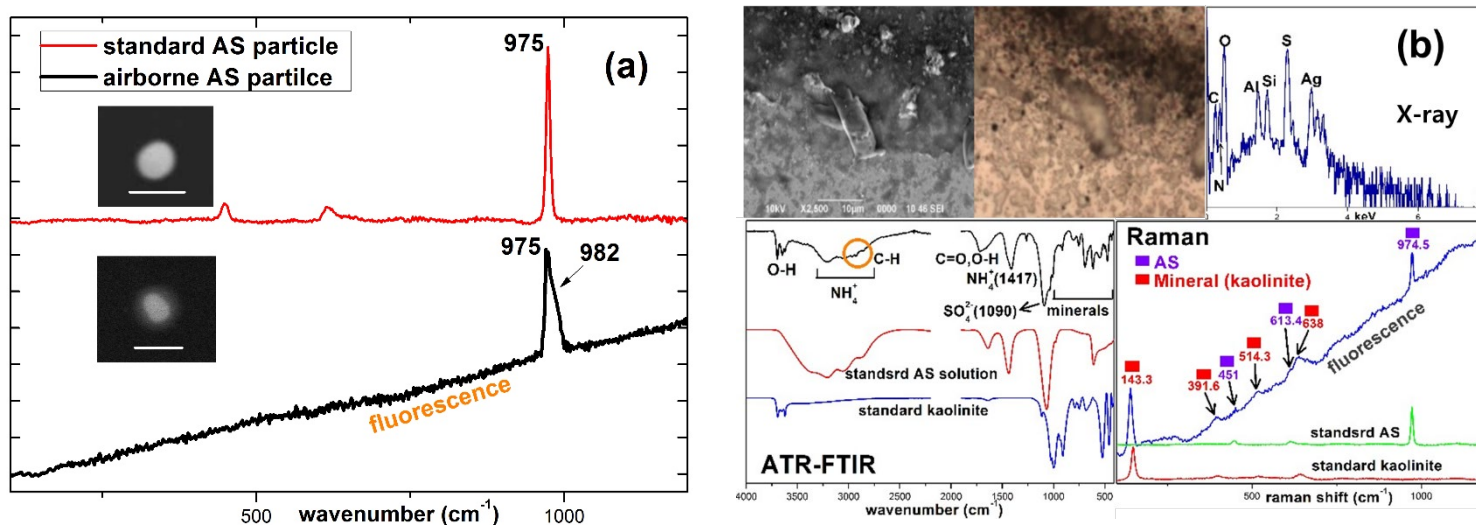


Figure 8. SEIs, X-ray spectra, and element atomic concentrations of aged mineral dust particles.

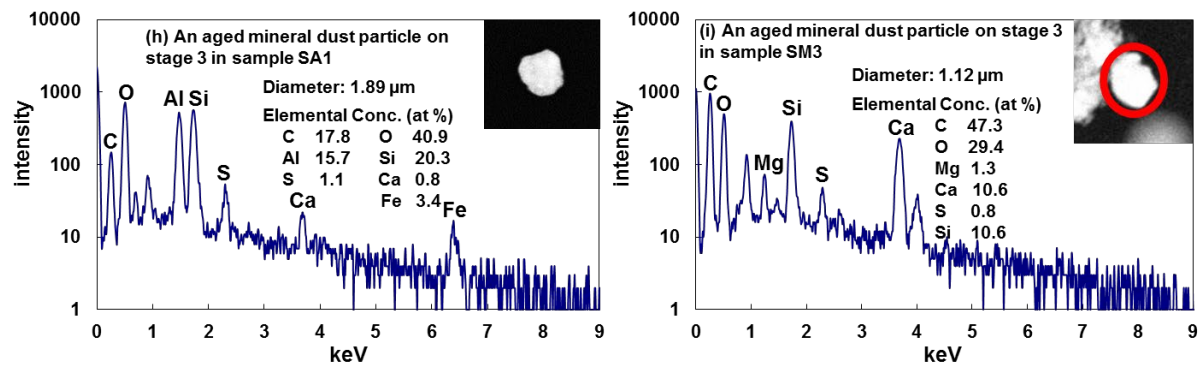
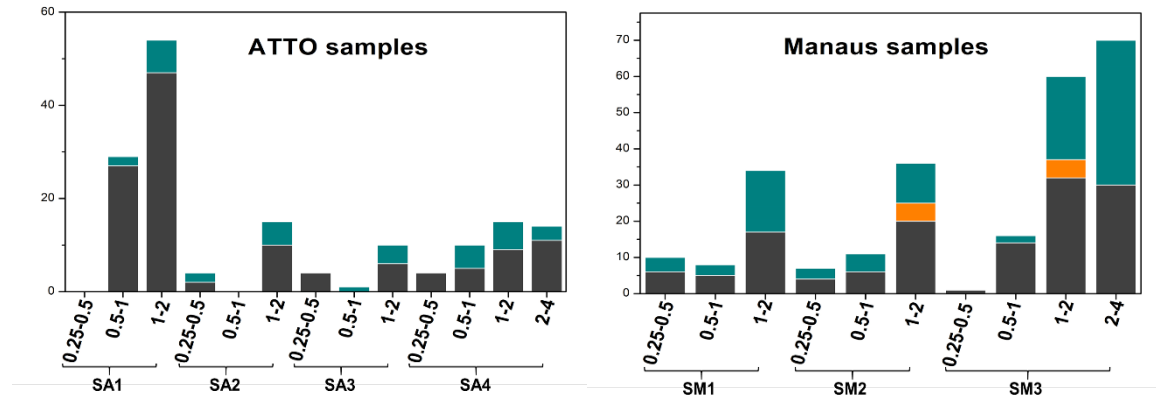


Figure 9. Number abundances of reacted sea-salt and aged mineral dust particles containing sulfates (■), nitrates (■), and both (■). The size ranges of each sample is given in the x-axis.

(a) aged mineral dust



(b) reacted sea-salts

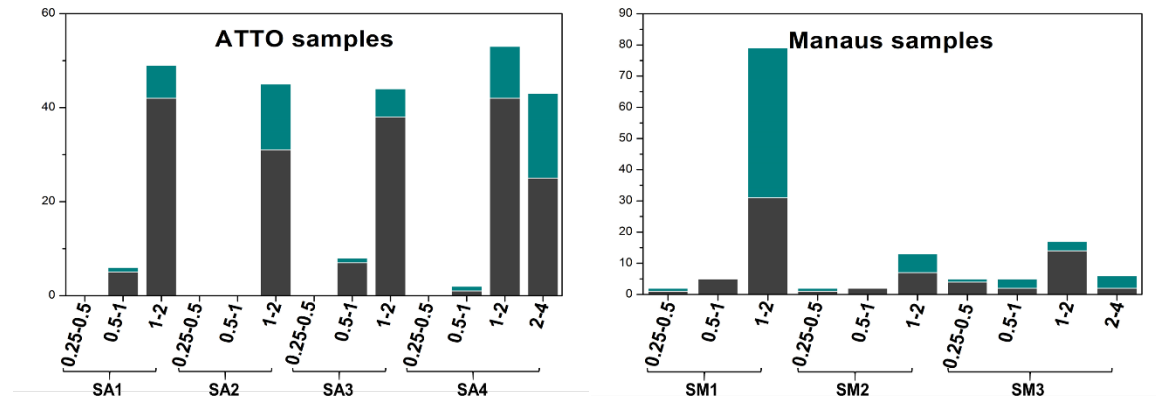


Figure 10. SEIs, X-ray spectra, and element atomic concentrations of (a) reacted sea-salt, (b) reacted sea-salt with K-salt, and (c) elongated CaSO_4

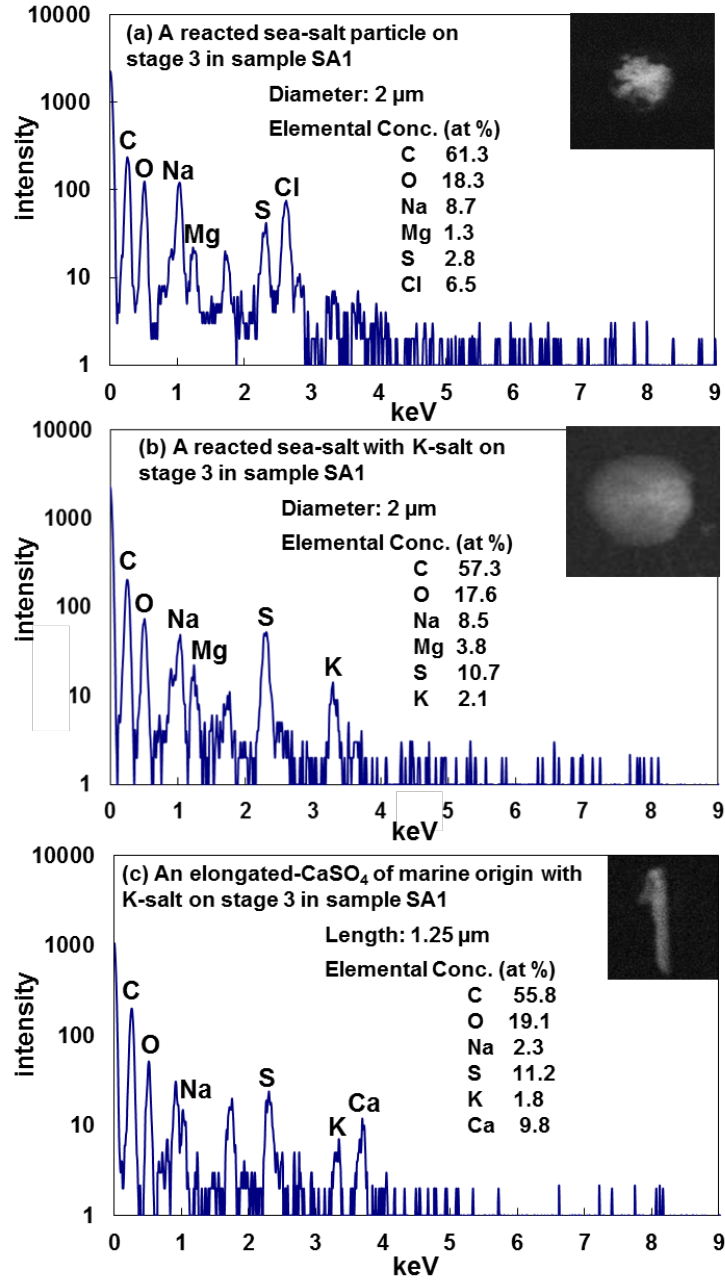


Figure 11. SEIs, X-ray spectra, and element atomic concentrations of (a) PBA, (b) soot, (c) tarball, (d) fly ash, (e) Ni-containing, and (f) Fe-containing particles.

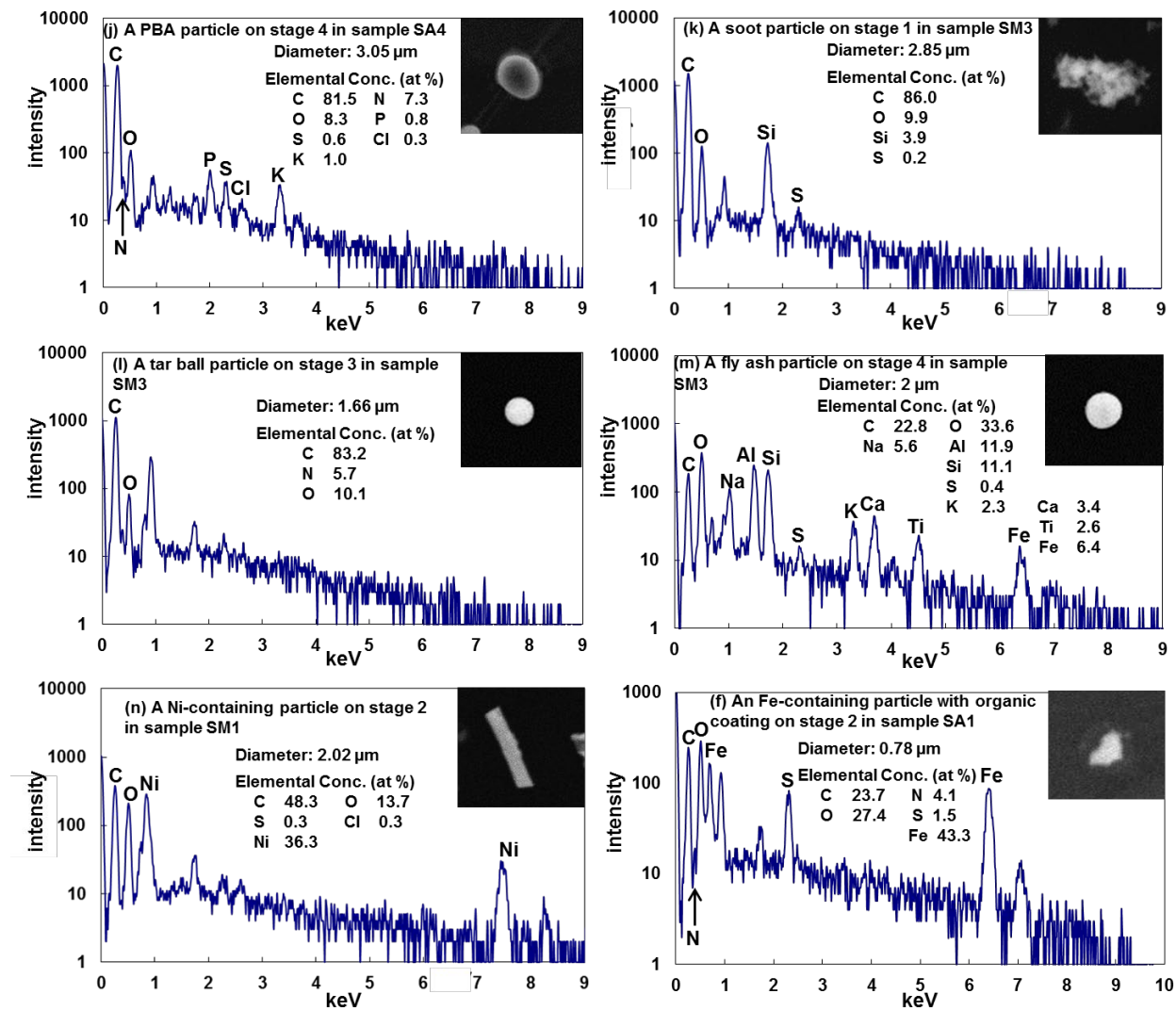


Figure 12. Typical SEIs of PBA particles from stage 4 of the (a) SA4 and (b) SM3 samples. PBA and PBA/SOA mixture particles are marked with (\rightarrow) and (+), respectively.

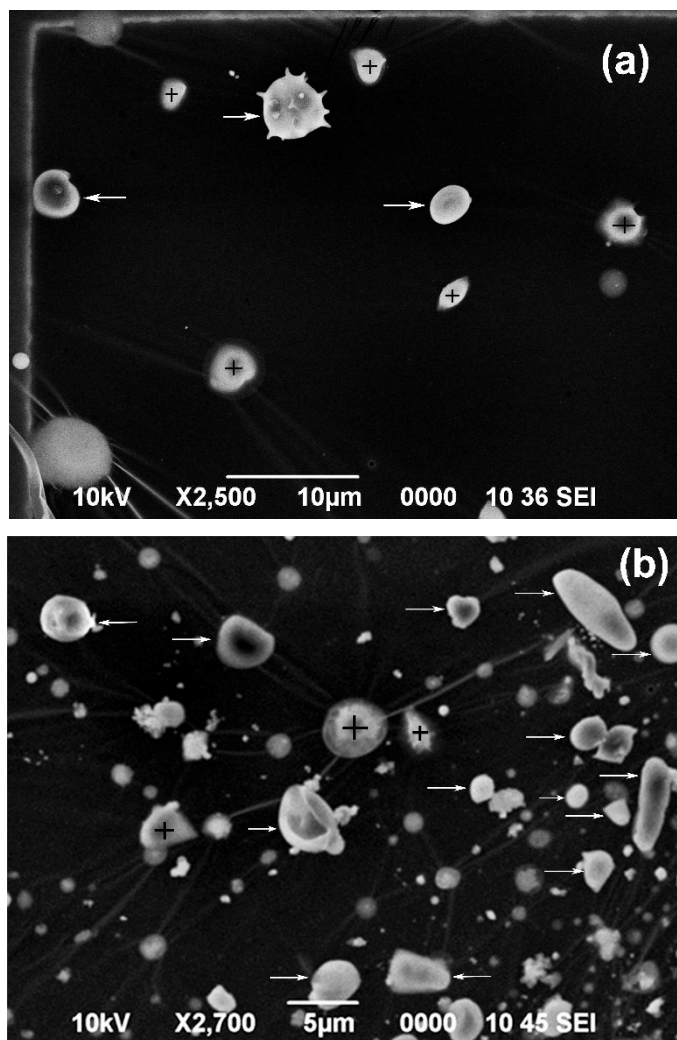


Figure 13. Relative number abundance of nine different particle types for the SA1-SA4 and SM1-SM3 samples collected at the ATTO and Manaus sites, respectively.

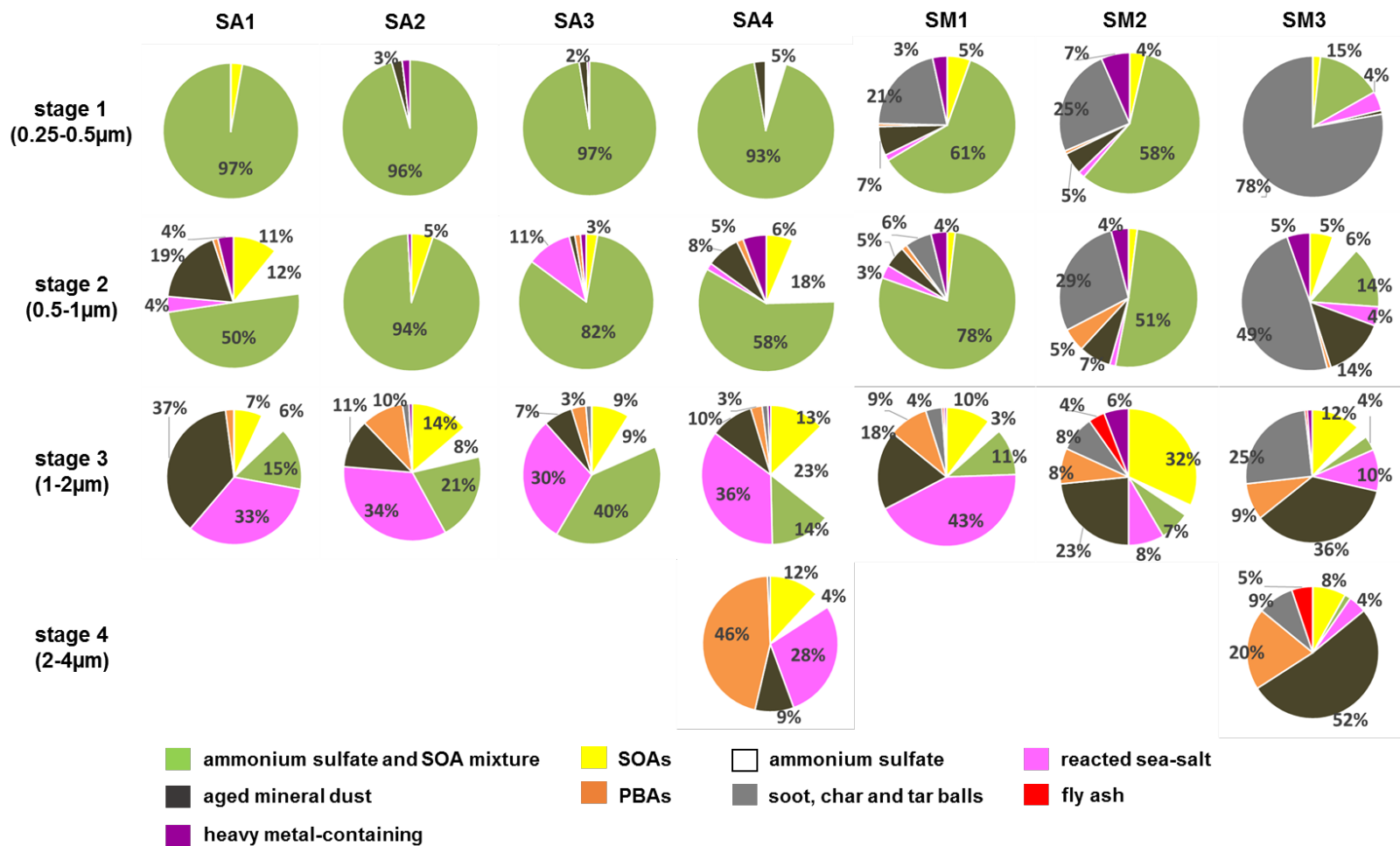


Table 1. Sampling dates, sampling times, and metrological conditions during the samplings.

Sampling site	Sample	Sampling date	Sampling time	<i>T</i> (°C)	RH (%)	Rainfall (mm)	Sampling conditions
ATTO	SA1	April 1, 2012	12:25 - 14:25	28	72	9	Precipitation during the TEM grid exposure
	SA2	April 16, 2012	12:12 - 13:52	28	74	0	Precipitation (1 mm) in the last 24 hours before the TEM grid exposure
	SA3	April 17, 2012	12:13 - 13:53	30	60	0	No precipitation
	SA4	April 18, 2012	12:04 - 13:45	27	75	0	No precipitation
Manaus	SM1	May 1, 2012	12:17 - 13:57	32	57	0	Precipitation (1 mm) in the last 24 hours before the TEM grid exposure
	SM2	May 2, 2012	12:00 - 13:40	27	73	0	Precipitation (2 mm) in the last 24 hours before the TEM grid exposure
	SM3	May 3, 2012	12:06 - 13:46	24	88	8	Precipitation during the TEM grid exposure