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**Title:** Analysis of New Particle Formation (NPF) Events at Nearby Rural, Urban Background and Urban Roadside Sites

**Author(s):** Dimitrios Bousiotis et al.

**MS Type:** Research article

**Iteration:** Minor Revision

## **RESPONSE TO THE CO-EDITOR**

### **Comments to the Author:**

Dear authors,

Thanks for revising the manuscript.

The main concern of the referees is the large lowest detectable size of 16.6 nm of the SMPS system. The referees are right that it is difficult to assess the aerosol formation with the measurements starting with such a large size. As a result, you have modified the manuscript and softened some of the conclusions. I still recommend to add a note about the measurement size also to the abstract and the conclusions.

**RESPONSE:** A note has been added in both the Abstract (line 31) and Conclusions (line 549) about the size range of the available data.

I have few additional comments that require some work.

Line 51: Clarify what do you mean by urban pollution.

**RESPONSE:** A clarification has been added about sources of urban pollution at the end of the abstract (line 51 of tracked version).

Line 125: What do you mean that particle formation was found to take place on event and non-event days.

**RESPONSE:** A clarification has been added about particle formation taking place but not qualifying as NPF events due to either newly formed particles not surviving or lack of growth (line 126).

Line 183: Confidence level typically implies statistical methods. Please modify the sentence.

**RESPONSE:** Confidence level has been changed to level of certainty (line 187)

Line 188: how frequent were the bursts detected with the CPC?

**RESPONSE:** An estimation of the frequency of the bursts and an explanation has been added (line 193). This estimation is based on a quick review of the data. To provide a more precise answer the whole dataset would need to be analysed again.

Line 223: What would the result be, if you would follow the mode through the night? Was growth typically persistent through the night? Setting a deadline for the midnight is rather arbitrary.

**RESPONSE:** A justification has been added of the reason why the end of the day was chosen as the final point for the growth rate calculation (line 228).

Line 237: NSF was proposed by who? Please add a reference.

**RESPONSE:** Reference added (line 245).

Line 280: From where was ammonia data available from? Please add a short section on ancillary data to the methods section.

**RESPONSE:** The data source has been added (line 166).

Line 299: From where did you get the organic compounds? VOC data it seems, but this is not described. Please describe in the methods section.

**RESPONSE:** The data source and measurement method has been added (line 168).

Line 315: Cluster 3 and following discussion is difficult to follow, but it clear after reading the section 3.3. Please summarise the trajectory analysis results here or consider structural changes.

**RESPONSE:** The chapter with the back trajectory analysis for the NPF events at the background sites was moved after the general back trajectory analysis for all three sites. Text and figure numbering were updated to match the changes made.

Line 320: again organic carbon concentration. From where?

**RESPONSE:** The data source has been added (line 169). Also, the text has been updated to indicate that the organic carbon concentration refers to both sites (line 487).

Line 323: low growth rate and consequently low survivability

**RESPONSE:** The text has been updated to address the suggested correction (line 490).

Line 329: How did you determine particulate organic carbon concentration?

**RESPONSE:** A data source has been added (line 167).

line 341: With the instrument that detects > 16.6. nm size distribution, one cannot assess the initial states of newly formed particles.

**RESPONSE:** The text has been updated to address the correction. “Initial stages” was changed to “early stages” (line 347).

Line 358: Ethane, from where is the data from?

**RESPONSE:** A data source has been added (line 167).

Table 1: A fraction of NPF days to all days would help to address the frequency at different locations.

**RESPONSE:** The number of days available per year were added in parentheses next to the number of events.

Table 2: What is the variability of these numbers?

**RESPONSE:** The table has been updated to present the variability of the values.

Figure 3: What is the variability of GR?

**RESPONSE:** The figure has been updated to present the variability of the values.

Figure 4: Mean or median? If the latter, quartile range would help to address the variability.

**RESPONSE:** The text of the figure has been updated to explain what is presented. The word “average” was replaced by the word “mean”.

Figure 6: Same comment as Figure 3.

**RESPONSE:** The figure has been updated to present the variability of the values.

Figure 7: same as above.

**RESPONSE:** The figure has been updated to present the variability of the values.

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3 **Analysis of New Particle Formation (NPF) Events at**  
4 **Nearby Rural, Urban Background and**  
5 **Urban Roadside Sites**  
6

7  
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## 29 ABSTRACT

30 New Particle Formation (NPF) events have different patterns of development depending on the  
31 conditions of the area in which they occur. In this study, [particle size distributions in the range of 16.6](#)  
32 [– 604 nm \(seven years of data\) were analysed and](#) NPF events occurring at three sites of differing  
33 characteristics (rural Harwell (HAR), urban background North Kensington (NK), urban roadside  
34 Marylebone Road (MR), London, UK) were [extracted and](#) studied ~~(seven years of data)~~. The different  
35 atmospheric conditions in each study area not only have an effect on the frequency of the events, but  
36 also affect their development. The frequency of NPF events is similar at the rural and urban  
37 background locations (about 7% of days), with a high proportion of events occurring at both sites on  
38 the same day (45%). The frequency of NPF events at the urban roadside site is slightly less (6% of  
39 days), and higher particle growth rates (average 5.5 nm h<sup>-1</sup> at MR compared to 3.4 nm h<sup>-1</sup> and 4.2 nm  
40 h<sup>-1</sup> at HAR and NK respectively) must result from rapid gas to particle conversion of traffic-generated  
41 pollutants. A general pattern is found in which the condensation sink increases with the degree of  
42 pollution of the site, but this is counteracted by increased particle growth rates at the more polluted  
43 location. A key finding of this study is that the role of the urban environment leads to an increment  
44 of 20% in N<sub>16-20nm</sub> in the urban background compared to that of the rural area in NPF events occurring  
45 at both sites. The relationship of the origin of incoming air masses is also considered and an  
46 association of regional events with cleaner air masses is found. Due to lower availability of  
47 condensable species, NPF events that are associated with cleaner atmospheric conditions have lower  
48 growth rates of the newly formed particles. The decisive effect of the condensation sink in the

49 development of NPF events and the survivability of the newly formed particles is underlined, and  
50 influences the overall contribution of NPF events to the number of ultrafine particles in an area. The  
51 other key factor identified by this study is the important role that ~~urban~~ pollution, both from traffic  
52 and other sources in the urban environment (such as heating or cooking), plays in new particle  
53 formation events.

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55

## 1. INTRODUCTION

Ultrafine particles (particles with diameter smaller than 100 nm) typically make the greatest contribution in the total particle count, especially in urban environments (Németh et al., 2018), but a very small contribution to total volume and mass (Harrison et al., 2000). Research studies have indicated that ultrafine particles can cause pulmonary inflammation and may contribute to cardiovascular disease (Oberdörster, 2000) and have increased possibility to penetrate the brain and central nervous system (Politis et al., 2008) compared to fine and coarser particles. Since some studies report that toxicity per unit mass increases as particle size decreases (Penttinen et al., 2001; MacNee et al., 2003; Davidson et al., 2005); it is considered possible that particle number concentrations may be a better predictor of health effects than mass concentrations (Harrison et al., 2000; Atkinson et al., 2010; Kelly et al., 2012; Samoli et al., 2016). Additionally, NPF events have an impact on climate (Makkonen et al., 2012) either by increasing the number of cloud condensation nuclei (Spracklen et al., 2008; Merikanto et al., 2009; Dameto de España et al., 2017; Kalkavouras et al., 2017), or directly affecting the optical properties of the atmosphere (Seinfeld and Pandis, 2012).

The sources of ultrafine particles in urban areas can either be primary particles or emission sources from traffic (Shi et al., 1999; Harrison et al., 2000), airports (Masiol et al., 2017) and other combustion related processes (Keuken et al., 2015; Kecorius et al., 2016), or by new particle formation (NPF) from gaseous precursors. New particle formation as described by Kulmala et al. (2014), is the process of production of low-volatility vapours, clustering of these vapours, nucleation, activation of the

76 clusters with a second group of vapours and condensational growth to larger sizes. This process can  
77 occur both locally or on a larger scale; in the latter case the events are characterized as regional.  
78 Regional events have been found to take place in a scale of hundreds of kilometres (Németh and  
79 Salma, 2014; Shen et al., 2018), without being affected by air mass advection (Salma et al., 2016).  
80 NPF is one of the main contributors of particles in the atmosphere (Spracklen et al., 2010; Kulmala  
81 et al., 2016; Rahman et al., 2017) and this relative contribution increases moving from a kerbside to  
82 a rural area (Ma and Birmili, 2015). While NPF events in rural and remote areas have been widely  
83 studied for many years (O'Dowd et al., 2002; Dal Maso et al., 2005; Ehn et al., 2010; Dall'Osto et al.,  
84 2017; Kalkavouras et al., 2017), in urban areas intensive studies have started mainly in recent years  
85 (Jeong et al., 2010; Minguillón et al., 2015; Peng et al., 2017; Németh et al., 2018). Early studies in  
86 Birmingham, UK highlighted the connection of NPF events with solar radiation (Shi et al., 2001) and  
87 a low condensation sink (Alam et al., 2003), a measure of pre-existing aerosol loading (Dal Maso et  
88 al., 2002). The importance of a low condensation sink was further underlined by later studies, as being  
89 one of the most influential variables in the occurrence of NPF in all types of environment (Wehner et  
90 al., 2007; Park, Yum and Kim, 2015; Pikridas et al., 2015). An important contributor to many NPF  
91 pathways is SO<sub>2</sub> (Woo et al., 2001; Berndt et al., 2006; Laaksonen et al., 2008), which in the presence  
92 of solar radiation forms H<sub>2</sub>SO<sub>4</sub>, often the main component of the initial clusters (Kuang et al., 2008;  
93 Kulmala et al., 2013; Bianchi et al., 2016; Kirkby et al., 2016). Dall'Osto et al. (2013) pointed out  
94 that the role of SO<sub>2</sub> is less significant in urban areas compared to rural and background areas. SO<sub>2</sub>  
95 concentration variability in urban areas was found to have a small impact on the frequency of NPF

96 events (Alam et al., 2003; Jeong et al., 2010), though it can have an effect on the number of particles  
97 formed (Charron et al., 2007). Furthermore, Dall'Osto et al. (2018) in their research at 24 sites in  
98 Europe, pointed out the different role SO<sub>2</sub> seems to play depending on its concentration, and that of  
99 other species. Jayaratne et al. (2017) however found that in the heavily polluted environment of  
100 Beijing, China, NPF events were more probable in sulphur rich conditions rather than sulphur poor.  
101 Apart from its role in the initial formation of the clusters, H<sub>2</sub>SO<sub>4</sub> seems to participate in the early  
102 stages of growth of the newly formed clusters (Kulmala et al., 2005; Iida et al., 2008; Xiao et al.,  
103 2015). In later stages of growth, low or extremely low volatility organic compounds (O'Dowd et al.,  
104 2002; Laaksonen et al., 2008; Metzger et al., 2010; Kulmala et al., 2013; Tröstl et al., 2016; Dall'Osto  
105 et al., 2018) appear to be more important, while the role of ammonium nitrate in particle growth is  
106 also considered (Zhang et al., 2017). While in rural areas the organic compounds are mainly of  
107 biogenic origin (Riccobono et al., 2014; Kirkby et al., 2016), in urban areas they mainly originate  
108 from combustion processes (Robinson et al., 2007; Gentner et al., 2012). Many comparative studies  
109 have reported higher growth rates in urban areas compared to background sites (Wehner et al., 2007;  
110 Jeong et al., 2010; Salma, et al., 2016; Wang et al., 2017), as well as greater particle formation rates  
111 (Salma, et al., 2016; Nieminen et al., 2018) and a higher frequency of NPF events (Peng et al., 2017),  
112 which was attributed to the higher concentration of condensable species. Salma et al. (2014) however  
113 reported fewer NPF events in the city centre of Budapest compared to the urban background, due to  
114 the higher condensation sink. Due to the complexity of the conditions and mechanisms within an  
115 urban area (Harrison, 2017), NPF events are harder to study and factors to be attributed. Increased

116 concentrations of particles in the size range 1.3 – 3 nm were measured at a kerbside site when  
117 downwind from the road, following the trends in traffic-related nucleation mode particles, associating  
118 them with traffic emissions and thus not resulting from homogeneous nucleation mechanisms  
119 (Rönkkö et al., 2017; Hietikko et al., 2018), and studies in Barcelona, Spain (Dall’Osto et al., 2012;  
120 Brines et al., 2014) and Leicester, U.K. (Hama et al., 2017), attributed a larger portion of nucleation  
121 mode particles to vehicular emissions compared to photochemically induced nucleation. As the  
122 condensation sink is higher within an urban environment, NPF events are less favoured. Their  
123 occurrence is attributed to either ineffective scavenging or the higher growth rate of the newly formed  
124 particles (Kulmala et al., 2017), when sufficient concentrations of precursors are present in the  
125 atmosphere (Fiedler et al., 2005), as particle formation was found to take place on both event and  
126 non-event days [with variable intensity, though not always followed by survival or growth of the newly](#)  
127 [formed particles, thus not qualifying as NPF events](#) (Riipinen et al., 2007).

128  
129 In this study, NPF events in three areas of different land use in the southern U.K. are analyzed. Studies  
130 for NPF events have been conducted in the past for Harwell, Oxfordshire (Charron et al., 2007; 2008)  
131 and the effect of NPF upon particle size distributions was also considered for N. Kensington, London  
132 (Beddows et al., 2015). A combined study including all three sites has also been conducted, but in  
133 the aspect of ultrafine particle variation (Von Bismarck-Osten et al., 2013). The present study is the  
134 first to use a combined long term database for all three sites, focusing on the trends and conditions of  
135 NPF events at these sites, as well as the first which identifies NPF events at the highly trafficked

Marylebone Road site, as up to this point ultrafine particles were attributed only to traffic (Charron and Harrison, 2003; Dall'Osto et al., 2011). As in this study a rural and an urban background area are studied alongside a kerbside site in the city of London in close proximity, the conditions and development of NPF events in a mid-latitude European region are discussed in relation to the influence of different local environments.

141

## 142 **2. DATA AND METHODS**

### 143 **2.1 Site Description and Data Availability**

144 This study analysed NPF events in three areas in the southern United Kingdom (Fig. 1). Harwell in  
145 Oxfordshire, is located about 80 km west of the greater London area. The site is in the grounds of the  
146 Harwell Science Centre in Oxfordshire (51° 34' 15" N, 1° 19' 31" W) and is representative of a rural  
147 background area; a detailed description of the site was given by Charron et al. (2013). North  
148 Kensington is a suburban area in the western side of London, U.K, 4.5 km west of Marylebone Road.  
149 The site is located in the grounds of Sion Manning School (51° 31' 15" N, 0° 12' 48" W) and is  
150 representative of the urban background of London. A detailed description of the site was given by  
151 Bigi and Harrison (2010). Marylebone Road is located in the centre of London, U.K. The site is  
152 located on the kerbside of Marylebone road (51° 31' 21" N; 0° 9' 16" W), a very busy arterial route  
153 within a street canyon. A more detailed description of the area can be found in Charron and Harrison  
154 (2003).

155 At all three sites, seven years (2009 – 2015) of particle number size distributions in the range of 16.6  
156 – 604 nm have been measured and recorded as 15-minute averages, using a Scanning Mobility  
157 Particle Sizer (SMPS), comprised by an Electrostatic Classifier (EC, TSI model 3080) and a  
158 condensation Particle Counter (CPC, TSI Model 3775), operated on behalf of the Department for  
159 Environment, Food and Rural Affairs (DEFRA) in the U.K. At all sites the inlet air is dried, and  
160 operation is in accord with the EUSAAR/ACTRIS protocol (Wiedensohler et al., 2012). These 15-  
161 minute measurements were averaged to an hourly resolution. In Harwell there were 46930 hours of  
162 available SMPS data (76.5% coverage), in N. Kensington 51059 (83.3% coverage) and at Marylebone  
163 Road 45562 (74.3% coverage). Detailed data availability is found in Table S1. A free-standing CPC  
164 (TSI model 3022A) also operated alongside for most of the years of the survey and was used to give  
165 an estimate of particles in the 7-16.6 nm range by difference from the SMPS.

166  
167 Additionally, air pollutants and other ~~aerosol-gas and particulate~~ chemical composition data (NO<sub>x</sub>,  
168 SO<sub>2</sub>, SO<sub>4</sub><sup>2-</sup>, Cl, Na, Mg, gaseous ammonia and VOCs) were extracted from the DEFRA website  
169 (<https://uk-air.defra.gov.uk/>); Daily measurements of particulate OC were also extracted from the  
170 DEFRA website which are determined- using the method described in the Annual report of the  
171 National Physicals Laboratory (Beccaceci et al., 2015). Meteorological data for Harwell and  
172 Heathrow airport (used for N. Kensington and Marylebone road) were available from the Met Office,  
173 while solar radiation data from Benson station (for Harwell) and Heathrow airport (for N. Kensington  
174 and Marylebone Road), were extracted from the Centre for Environmental Data Analysis (CEDA)

175 site (<http://www.ceda.ac.uk>). Back trajectory data calculated using the HYSPLIT model (Draxler and  
176 Hess, 1998), were extracted by the NOAA Air Resources Laboratory  
177 (<https://ready.arl.noaa.gov/READYtransp.php>) and were processed using the Openair package for R  
178 (Carslaw and Ropkins, 2012).

179

## 180 **2.2 Methods**

### 181 **2.2.1 NPF events selection**

182 The identification of the NPF event days was made by visual inspection of SMPS data, supplemented  
183 with the use of CPC data to confirm the formation of a new mode of particles, using the criteria set  
184 by Dal Maso et al. (2005). NPF events are considered when a distinctly new mode of particles which  
185 appears in the size distribution at nucleation mode size, prevails for some hours and shows signs of  
186 growth. Using these criteria, NPF events are classified into two classes, I and II depending on the  
187 ~~confidence-level~~ [of certainty](#). Class I events are further classified to Ia and Ib, with class Ia containing  
188 very clear and strong particle formation events, while Ib contains less clear events. In this study the  
189 events of class Ia only are considered as being the most suitable for analysing case studies of NPF  
190 events (Figure S1). At this point it should be mentioned that due to the particle size range available,  
191 NPF events in which new formed particles failed to grow beyond 16.6 nm (if any) could not be  
192 identified. ~~Though such rare occasions were identified using the CPC data, b~~Bursts of new particles  
193 in the size range < 16.6 nm ~~that were identified using the CPC data -that~~but did not appear in the  
194 SMPS dataset were ignored as their development was unknown. [This type of development was rare](#)

195 [and mainly found ~~in~~at the rural background site, occurring on a few days per year mainly in summer.](#)  
196 [Its main feature was the short duration of the bursts compared to event days. In the urban sites, this](#)  
197 [type of development was almost non-existent.](#)- High time resolution data for gaseous pollutants and  
198 aerosol constituents was used to identify pollution events affecting particle concentrations and these  
199 were removed from the data analysis. This analysis took account of the fact that nanoparticle  
200 emissions from Heathrow Airport affect size distributions at London sites (Harrison et al., 2018), and  
201 such primary emission influences were not included as NPF events.

202

### 203 **2.2.2 Calculation of the condensation sink and growth rate**

204 For the calculation of the condensation sink the method proposed in Kulmala et al. (2001) was used  
205 in which the condensation sink is calculated as

206

$$207 \quad CS = 4\pi D \sum \beta_M r N \quad (1)$$

208

209 where  $r$  is the radius of the particles and  $N$  is the number concentration of the particles.  $D$  is the  
210 diffusion coefficient calculated (for  $T = 293$  K and  $P = 1013.25$  mbar) according to Polling et al.  
211 (2000):

212

$$D_{\text{vap}} = 0.00143 \cdot T^{1.75} \frac{\sqrt{M_{\text{air}}^{-1} + M_{\text{vap}}^{-1}}}{P \left( D_{\text{x,air}}^{\frac{1}{3}} + D_{\text{x,vap}}^{\frac{1}{3}} \right)^2} \quad (2)$$

where P is air pressure, M is the molar mass and  $D_x$  is the diffusion volume for air and  $\text{H}_2\text{SO}_4$ .  $\beta_M$  is the Fuchs correction factor calculated as (Fuchs et al., 1971):

$$\beta_M = \frac{1 + K_n}{1 + \left( \frac{4}{3a} + 0.377 \right) K_n + \frac{4}{3a} K_n^2} \quad (3)$$

where  $K_n$  is the relation of the particle diameter and the mean free path of the gas  $\lambda_m$ , called the Knudsen number.

The growth rate of the particles on nucleation event days was also calculated as proposed by Kulmala et al. 2012, using the formula

$$GR = \frac{D_{p_2} - D_{p_1}}{t_2 - t_1} \quad (4)$$

for the size range 16.6 – 50 nm. The number of points taken depended on the development of the event and were considered from the start of the event until a) growth stopped, b) GMD reached 50

230 nm or c) the day ended [\(this cut-off was chosen as the development of an event in its later stages is](#)  
231 [heavily biased by the local conditions, especially ~~in~~at the urban sites\).](#)

232

### 233 2.2.3 Calculation of the urban increment (U.I.)

234 The urban increment is defined as the ratio of the number concentration of particles below 20 nm for  
235 event days to the average (for the period April – October, when the majority of the events take place)  
236 for North Kensington to that at Harwell. This provides with a measure of the new particles formed in  
237 each area in comparison to the average conditions, and is calculated by

$$238 \text{ U.I.} = \frac{NK_{\text{Nuc Max}} - NK_{\text{Bg}}}{HW_{\text{Nuc Max}} - HW_{\text{Bg}}} \quad (5)$$

239

240 where  $NK_{\text{Nuc Max}}$  is the maximum concentration of particles below 20 nm found in the diurnal cycle  
241 on event days (found at 13:00) and  $NK_{\text{Bg}}$  is the average mean concentration at the same time (same  
242 for Harwell in the denominator).

243

### 244 2.2.4 Calculation of nucleation strength factor (NSF) and the P parameter

245 The Nucleation Strength Factor (NSF) was proposed [by Salma et al. \(2014\)](#) as a measure of the effect  
246 nucleation events have in the composition of ultrafine particles in an area. Two factors were proposed.

247 First is the  $NSF_{\text{Nuc}}$ . This is calculated as

248

$$NSF_{\text{NUC}} = \frac{\left( \frac{N_{(\text{smallest size available}-100)}}{N_{(100-\text{largest size available})}} \right)_{\text{nucleation days}}}{\left( \frac{N_{(\text{smallest size available}-100)}}{N_{(100-\text{largest size available})}} \right)_{\text{non-nucleation days}}} \quad (6)$$

249

250

251 and provides of a measure of the concentration increment on nucleation days exclusively caused by

252 new particle formation (NPF). The second factor is  $NSF_{\text{GEN}}$  calculated as

253

$$NSF_{\text{GEN}} = \frac{\left( \frac{N_{\text{smallest size available}-100}}{N_{100-\text{largest size available}}} \right)_{\text{all days}}}{\left( \frac{N_{\text{smallest size available}-100}}{N_{100-\text{largest size available}}} \right)_{\text{non-nucleation days}}} \quad (7)$$

254

255 and gives a measure of the overall contribution of NPF on a longer span (Salma et al. 2017).

256 The dimensionless survival parameter P, as proposed in Kulmala et al. (2017), was calculated as

257

$$P = \frac{CS'}{GR'}$$

258

259 where  $CS' = CS/(10^{-4} \text{ s}^{-1})$  and  $GR' = GR/(1 \text{ nm hour}^{-1})$ . CS and GR values used were calculated with

260 the methods mentioned at 2.2.2. An increased P parameter is an indication that a smaller percentage

261 of newly formed particles will survive to greater sizes. Hence this is the inverse of particle

262 survivability, and values of  $P < 50$  are typically required for NPF in clean or moderately polluted

environments, although higher values of P are observed in highly polluted atmospheres (Kulmala et al, 2017).

### **3. RESULTS AND DISCUSSION**

#### **3.1 NPF Events in the Background Areas**

##### **3.1.1 Conditions and trends of NPF events**

The number of NPF event days for each site per year, those that took place simultaneously at both urban and rural background sites, as well as those events that took place at all three sites simultaneously appear in Table 1. Given that overall data recovery was in the range of 74-83%, results from individual years are unreliable, but the seven-year runs should average out most of the effects of incomplete data recovery. The number of events is similar for Harwell and N. Kensington, with a frequency of about 7% of all days with data. There is a clear seasonal variation favouring summer and spring (Figure 2) for both areas of the study. A similar pattern of variation was found for N. Kensington by Beddows et al. (2015). In general, higher solar radiation, lower relative humidity, low cloud cover and higher pressure conditions, lower concentrations of pollutants as well as lower condensation sink are found when NPF events took place in all areas (Figure S2), as was also reported by Charron et al. (2007) for Harwell. While SO<sub>2</sub> is one of the main factors for NPF events to occur, concentrations are lower when events take place. This is indicative that SO<sub>2</sub> concentrations in these areas are sufficient for events to take place, and higher concentrations are likely to be associated with higher pollution and a higher condensation sink. The proxy for [H<sub>2</sub>SO<sub>4</sub>] was calculated for the

background sites using the method outlined in Petäjä et al., (2009) and was found to be higher on event days for both background sites (results not included). This indicates the possible positive effect of increased concentrations of  $\text{H}_2\text{SO}_4$  in the occurrence of NPF events as well as, since  $\text{SO}_2$  concentrations were found lower, the increased role of either the solar radiation (via the formation of OH radical) or the reduced condensation sink to its formation. For the case of gaseous ammonia (results not included) for Harwell where data was available, as there was no distinct variation found between event and non-event days, but as the concentration of ammonia in the U.K. is in the range of few ppb (Sutton et al., 1995), it is sufficient according to ternary nucleation theory (Korhonen et al., 1999) for NPF events not to be limited by ammonia. The average growth rate for Harwell was found to be  $3.4 \text{ nm h}^{-1}$ , within the range given by Charron et al. (2007) and higher at N. Kensington at  $4.2 \text{ nm h}^{-1}$ , a trend found for all seasons (Figure 3). The increased growth rate in the urban area can be related to the greater presence of organic matter and other condensable species. In both areas NPF events had higher growth rates in summer than in spring, as was also found in previous studies (Kulmala et al., 2004; Nieminen et al., 2018). This may be associated with the higher concentration of organic compounds emitted by trees during summer (Riipinen et al., 2007), or faster oxidation rates due to higher concentrations of hydroxyl radical and ozone (Harrison et al., 2006).

About 45% of the events took place simultaneously in both background areas. These events are characterized as regional, as NPF took place on a larger scale, regardless of the local conditions of the given area. In this case, meteorological conditions were even clearer, indicative of the greater

dependence of regional events on synoptic conditions rather than local. While most chemical constituents were also lower in concentration during regional events, different patterns were found for organic compounds and sulphate for each background area. In Harwell sulphate was higher during regional events, while in N. Kensington organic compounds were higher during regional events. This may be indicative of the variable role that specific chemical species have in condensational nanoparticle growth (Yue et al., 2010). In all cases though, the concentrations of these species were lower compared to the average conditions. Despite these differences, the growth rate of particles was found to be higher for local events in N. Kensington ( $4.4 \text{ nm h}^{-1}$ ) compared to regional events ( $3.9 \text{ nm h}^{-1}$ ), though within the margin of uncertainty. In Harwell, no difference was found in the growth rate between regional and local events.

313

### 314 **3.1.2 — Variability of the origin of the air masses on NPF events**

315 ~~As both background sites are relatively close to each other (about 80 km) and had similar number of~~  
316 ~~event days, a combined clustering of back trajectories for the event days (only) in these two areas was~~  
317 ~~attempted. This would provide an insight into the origin of air masses for local and regional events,~~  
318 ~~as well as the conditions for these air masses. The data for local N. Kensington events and both local~~  
319 ~~and regional events in Harwell were clustered together and the results along with the characteristics~~  
320 ~~of the air mass clusters are found in Figure S3.~~

321

Cluster C3, which is placed between C2 and C4 among those originating from the Atlantic Ocean, has the highest percentage for both area specific and regional events. Specifically, for regional events the percentage is over 35%, much higher compared to all other, showing a clear “preference” of regional events for cleaner and faster moving air masses from mid-latitudes of the Atlantic Ocean. This “preference” explains the lower production and growth rate of the new particles found for regional events, compared to local ones, as air masses from this area have lower organic carbon and SO<sub>2</sub> concentrations. Cluster C5, originating straight from the north but representing air masses that have crossed the Irish Sea and have not extensively gone over land presents a similar case. These cold and clean air masses are associated with a low growth rate and survivability of the newly formed particles. Local events for both sites apart from those in Cluster C3 are highly associated with Clusters C1 and C2. C1, which contains slow and polluted air masses, presents the highest growth rate and as a result high particle survivability, as given by the P parameter (see later). On the other hand, C2 which consists of warm and moist air masses from lower latitudes is the least common for regional events and presents high growth rate and survival probability of the particles. Apart from the weak relation found with particulate organic carbon concentrations and growth rate (Figure S3), there appears to be an inverse relation between the temperature and survivability of the particles. Warmer air masses seem to be related to higher particle survival probability, which may be attributable to greater growth rates as temperature increases (Yli-Juuti et al., 2011).

340

### 3.1.23 Urban increment and particle development

342 The urban environment, depending on the conditions, may have a positive or negative effect in the  
343 number of the particles formed and their consequent survival and growth. Both Harwell and N.  
344 Kensington are in background areas, rural and urban respectively. As a result, while the  
345 concentrations of pollutants are higher in N. Kensington than Harwell, their effect is smaller  
346 compared to that of Marylebone Road. A comparison of the particles smaller than 20 nm, gives insight  
347 into the formation and survival of the newly formed particles [in the early stages in the initial stages](#).  
348 Calculating the urban increment (equation 5) using the two background sites showed around 20%  
349 more particles of size 16 - 20 nm in N. Kensington than Harwell for event days, an increment that is  
350 even stronger when solely local events are considered (Figure 4). As the sizes of the particles in the  
351 calculation are relatively large and due to the higher condensation sink found in N. Kensington, this  
352 increment is expected to be larger for smaller size particles. A possible explanation for this result may  
353 be the greater concentration of organic compounds which is observed in N. Kensington, as discussed  
354 earlier, which leads to more rapid formation of secondary condensable species that enhances the  
355 nucleation process in the more polluted area.

356

357 Considering the local events, most of the pollutant concentration data available appear to be higher  
358 which is reflected in the condensation sink as well. The role of the polluted background appears to be  
359 decisive in the further growth of the newly formed particles, especially for Harwell. This, at both sites  
360 causes the number of particles of greater size to be smaller for the later hours in the days of local  
361 events (Figure S34). Another possible reason for this difference in the larger size ranges can be the

362 higher concentration of organic content on the days of regional events at N. Kensington (as discussed  
363 earlier). On the other hand, for Harwell all hydrocarbons with available data are lower throughout the  
364 day (apart from ethane) during regional events. Unlike N. Kensington, at Harwell particles smaller  
365 than 20 nm as well as the growth rate of the newly formed particles are almost the same for regional  
366 and local events.

367

368 The calculation of the increment in Marylebone Road provided negative results; particles smaller than  
369 20 nm were less abundant on event days compared to the average, throughout the day. This is due to  
370 the fact that Marylebone road is heavily affected by traffic pollution and on average, conditions do  
371 not promote NPF events due to the high condensation sink, unless clear conditions prevail, which are  
372 also associated with a low particle load.

373

### 374 **3.2 NPF Events at Marylebone Road**

375 For many years, NPF events were thought not to take place in heavily polluted urban areas, as the  
376 effect of the increased condensation sink was considered crucial in suppressing the formation and  
377 growth of new particles. Recent long term analyses have shown this is not the case and nowadays an  
378 increasing number of studies confirm the occurrence of NPF events in urban areas. In this study, for  
379 the same period of seven years as for the two background areas, NPF events were found to occur for  
380 6.1% of days at Marylebone Road, lower than in the background areas. Though, due to the particle  
381 size range available there cannot be a definitive answer to whether the formation of the particles takes

place in the specific locality of the sampling site, due to the observed increase in particle concentrations in the range 7 – 16 nm (provided by the CPC data) and the increased growth rates found in urban areas in general, it can be assumed that the formation takes place either in the area of the measuring site or in its close vicinity, while the growth of the particles persists in the area for several hours, despite the high condensation sink. Seasonal variation is similar to that at the background sites, but day of the week variation is stronger at Marylebone Road further favouring weekends (Figure S45), as on these days traffic intensity is lower.

In general, similar conditions found to affect NPF events at the background sites are also found at Marylebone Road, despite a much larger condensation sink. (Figure S2). As a result, less particles of size smaller than 20 nm were found on NPF event days than the average for the site, as the sum of background particles plus those formed on these days were less than that on an average day. The growth rate of the newly formed particles ( $5.5 \text{ nm h}^{-1}$ ), is higher than that of the background sites which is in agreement with the findings in the study of the background areas on the possible role of the condensable species, the concentrations of which are even greater at the urban kerbside. About 15% of NPF event days at Marylebone Road presented particle shrinkage after the initial growth; the study of these cases though is outside of the context of the present work. At Marylebone Road, the number of NPF days which were common with the background sites was fewer, as local conditions (high condensation sink) are detrimental to the occurrence of NPF events and thus the days of regional events including Marylebone Road were separately studied for this site. The regional event days that

were common for all three sites were 37 (31% of events at Marylebone Road) (Table 1). As with the other two areas, the growth rate is higher during local events, but the conditions are mixed, with lower concentrations of sulphate and organic compounds but higher SO<sub>2</sub>, NO<sub>x</sub> and elemental carbon. The relationship with higher wind speed (mainly western) (Figure S6), solar radiation (which results in greater H<sub>2</sub>SO<sub>4</sub> formation) and lower relative humidity, indicate the stronger relation of the regional events with synoptic conditions than the local events in the heavily polluted environment of Marylebone Road.

### 3.3 Connection of NPF Events with Incoming Air Masses

#### 3.3.1 Air mass back trajectory clustering and connection with NPF events

The origin of the air masses plays a very important role in the occurrence of NPF events, as shown in Section 3.1.2. Air masses of different origins have different characteristics. Back trajectories provide excellent insight into the source of the air masses. Air mass back trajectories were calculated both for all days and for NPF event days for each site separately, with the aim of complementing the analysis in Section 3.1.2 which addressed only the event days. The additional analysis gives a view of the frequency of NPF events within different air mass types. The initial air mass back trajectory clustering ended up with an optimal solution of 9 clusters of different air masses. As many of these clusters had similar characteristics and origin, solutions with fewer clusters were attempted. As the number of clusters was decreasing clusters became a mixture of different origins, thus making the distinction of different sources harder. As a result, the method chosen was to merge clusters of similar

422 origin and characteristics, which kept the detail of the large number of clusters and made the  
423 separation of the different origins more distinct.

424

425 The resulting four merged clusters (Figure 5), using the characterisation proposed by McIntosh et al.  
426 (1969) are:

- 427 • An **Arctic** cluster, which originates mainly from the northerly sector. It occurs about 10% of the  
428 time and consists of cold air masses, which either passed over northern parts of the U.K. or  
429 through the Irish Sea.
- 430 • A **Tropical** cluster, which originates from the central Atlantic. It occurs 25% of the time and  
431 contains warmer air masses. A small percentage of this cluster contains masses that have passed  
432 over countries south of the U.K. Even though these days were more polluted, the clustering  
433 method was unable to clearly distinguish these days as it does not take into account particle  
434 numbers or composition, even when the 9-cluster solution was applied.
- 435 • A **Polar** cluster, which originates from the north Atlantic. It is the most common type of air mass  
436 arriving in the areas of study and occurs about 40% of the time bringing fast moving, “clean” air  
437 masses with increased marine components (Cl, Na, Mg) from the west. This cluster also contains  
438 airmasses that have passed through Ireland, though an effect on particle size and chemical  
439 composition is not distinct.
- 440 • A **Continental** cluster, which originates from the east. It occurs about 25% of the time and  
441 consists mainly of slow moving air masses, originating from the London area (for the background

442 areas) and/or continental Europe. It has higher concentrations of most pollutants as well as the  
443 highest condensation sink.

444  
445 The occurrence of each air mass class for average and event days for Harwell and London (both sites)  
446 can also be found in Figure 5, while their main characteristics for each site can be found in Table S2.  
447 Though in this case the air mass grouping for each site was done in a different analysis, the resulting  
448 groups are almost identical in their characteristics and frequency, as the sites are close to each other.  
449 The Polar cluster is the one prevailing on both average and event days. This consists of clean fast-  
450 moving air masses originating mainly from mid and high latitudes of the Atlantic, and this cluster  
451 presents favourable conditions for NPF events. The association of NPF events with air masses from  
452 the mid-Atlantic at N. Kensington was also found by Beddows et al. (2015). Cool Arctic air masses  
453 on average are not clean as they may have passed over the northern U.K. The event days associated  
454 with this air mass type have the lowest concentrations of the pollutants within available data for all  
455 areas. The increased percentage of events with this air mass at all sites indicates that lower  
456 temperatures, in a clear atmosphere with sufficient solar radiation are favourable for NPF events as  
457 found in previous studies (Napari et al., 2002; Jeong et al., 2010; Kirkby et al., 2011). A similar trend  
458 of increased probability with polar and arctic maritime air masses was also found for Hyytiälä,  
459 Finland by Nilsson et al. (2001). Tropical air masses have a lower probability for NPF events, which  
460 is associated with the fact that a number of these days are associated with air masses which have  
461 passed from continental areas south of the U.K. (France, Spain etc.). Specifically for Marylebone

462 Road the NPF probability is a lot lower (11% versus 17% for N. Kensington and 20% for Harwell).  
463 This is due to the fact that these air masses are more related to southerly winds which, in Marylebone  
464 Road are associated with a street canyon vortex which causes higher pollutant concentrations at this  
465 site. Finally, the Continental cluster presents the lowest probability for NPF events. The air masses  
466 in this group originate from continental Europe and for the background areas in most cases have  
467 passed over the London region as well. This results in both a higher condensation sink and  
468 concentration of pollutants, which limits the number of days with favourable conditions for NPF  
469 events. Growth rate for all sites though appears to be higher for air masses originating from more  
470 polluted areas (Figure 6), which appear to enhance the growth process due to containing a higher  
471 concentration of condensable species (after oxidation).

472

### 473 **3.3.2 Variability of the origin of the air masses on NPF events**

474 As both background sites are relatively close to each other (about 80 km) and had similar number of  
475 event days, a combined clustering of back trajectories for the event days (only) in these two areas was  
476 attempted. This would provide an insight into the origin of air masses for local and regional events,  
477 as well as the conditions for these air masses. The data for local N. Kensington events and both local  
478 and regional events in Harwell were clustered together and the results along with the characteristics  
479 of the air mass clusters are found in Figure S5.

480

481 [Cluster C3, which is placed between C2 and C4 among those originating from the Atlantic Ocean,](#)  
482 [has the highest percentage for both area specific and regional events. Specifically, for regional events](#)  
483 [the percentage is over 35%, much higher compared to all other, showing a clear “preference” of](#)  
484 [regional events for cleaner and faster moving air masses from mid-latitudes of the Atlantic Ocean.](#)  
485 [This “preference” explains the lower production and growth rate of the new particles found for](#)  
486 [regional events, compared to local ones, as for air masses from this area lower organic carbon and](#)  
487 [SO<sub>2</sub> concentrations were found at both sites in this study. Cluster C5, originating straight from the](#)  
488 [north but representing air masses that have crossed the Irish Sea and have not extensively gone over](#)  
489 [land presents a similar case. These cold and clean air masses are associated with a low growth rate](#)  
490 [and consequently low survivability of the newly formed particles. Local events for both sites apart](#)  
491 [from those in Cluster C3 are highly associated with Clusters C1 and C2. C1, which contains slow and](#)  
492 [polluted air masses, presents the highest growth rate and as a result high particle survivability, as](#)  
493 [given by the P parameter \(see later\). On the other hand, C2 which consists of warm and moist air](#)  
494 [masses from lower latitudes is the least common for regional events and presents high growth rate](#)  
495 [and survival probability of the particles. Apart from the weak relation found with particulate organic](#)  
496 [carbon concentrations and growth rate \(Figure S5\), there appears to be an inverse relation between](#)  
497 [the temperature and survivability of the particles. Warmer air masses seem to be related to higher](#)  
498 [particle survival probability, which may be attributable to greater growth rates as temperature](#)  
499 [increases \(Yli-Juuti et al., 2011\).](#)

### 501 **3.4 Nucleation Strength Factor (NSF)**

502 The NSF (equations 6 and 7) is used to describe the effect nucleation events have on the number of  
503 particles at a site. The values of NSF for each site and for seasons spring and summer are shown in  
504 Table 2. The decrease of the contribution of NPF events to particle number, moving from the rural  
505 area to the kerbside was also found in previous studies (Salma et al., 2014; 2017). This is explained  
506 by the increased contribution to the particle number concentrations of other sources, mainly  
507 combustion in the urban environment, compared to rural areas. Apart from this trend, in the  
508 background areas the increase of  $N_{16-100}$  was greater in spring than summer. This effect seems stronger  
509 in the urban background area compared to the rural, as in that area the variability of  $N_{16-100}$  is greater  
510 for event days compared to that of the rural area. On the other hand, the contribution of NPF events  
511 in the longer span, as is illustrated by the  $NSF_{GEN}$  appears to favour summer for all areas, showing  
512 the increased formation and survivability of particles in this season.

513  
514 For Marylebone Road the result for the increase of the  $N_{16-100}$  is greater in summer than in spring, in  
515 contrast to what was found for the background sites. This is due to the fact that in summer the traffic  
516 intensity is decreased, giving the contribution from NPF events a stronger effect compared to the  
517 other sources. The very small increase found on NPF events in Marylebone Road, with a factor of  
518 just 1.26, a lot lower than that found in the urban area of Seoul, South Korea (Park et al., 2015), is  
519 indicative of the reduced effect of NPF events in an area which is heavily affected by traffic, as also

520 pointed out by Von Bismarck-Osten et al. (2013) in their study on particle composition in Marylebone  
521 Road.

522

### 523 **3.5 The Survival Parameter P**

524 The average values of the P parameter for each of the areas of this study are 10.5 for Harwell, 15.8  
525 for N. Kensington and 28.9 for Marylebone Road. The values found put Marylebone Road to the  
526 upper end of heavily polluted areas in Europe, North Kensington to the same level as many other  
527 urban areas in Europe, while Harwell had somehow higher values compared to other rural background  
528 areas in Europe, as calculated by Kulmala et al. (2017). The seasonal, air mass origin and local versus  
529 regional variations can be found in Figure 7 (winter is excluded due to very low number of events).  
530 While the increasing trend of the P parameter as we move from rural background to kerbside was  
531 expected, it can be seen that there is a clear seasonal pattern in all three areas, with summer having  
532 the lowest P parameter (greatest survivability) compared to the other two seasons. This is associated  
533 with the higher growth rate found in summer for all areas of this study, as the differences in the  
534 condensation sink on event days are negligible between seasons. The case is similar for regional and  
535 local events. The result per air mass origin is related to the different conditions and parameters of  
536 each incoming air mass in each area. For example, the higher P parameter for Tropical air masses at  
537 Marylebone Road, is associated with the higher condensation sink found for this kind of air masses,  
538 due to the street canyon effect which is specific for Marylebone Road for southerly wind directions  
539 with which these air masses are mainly related, while the higher values for the rather clean Arctic air

540 masses for the other two areas are associated with the lower growth rates found for this kind of air  
541 mass in these areas. The more polluted Continental air masses seem to have a different effect for rural  
542 and urban areas. Their higher condensation sinks and concentrations of pollutants have a negative  
543 effect on P-values for the rural site and a positive effect at the urban sites. The exact opposite is found  
544 for the cleaner air masses of the Polar cluster, which appear to result in reduced P-values of the newly  
545 formed particles at the urban sites. This is related to the lower condensation sink associated with this  
546 air mass type.

547

#### 548 **4. CONCLUSIONS**

549 Seven years of [particle size distributions in the range 16.6 – 604 nm and other meteorological and](#)  
550 [chemical composition](#) data from three distinct areas (regional background, urban background,  
551 kerbside) in the southern U.K. were analysed and the conditions associated with NPF events were  
552 studied. NPF events were found to occur on about 7% of days at background sites and less at the  
553 kerbside site. The conditions on event days for all three areas were similar, with clear atmospheric  
554 conditions and a lower condensation sink. While the condensation sink appears to be the most  
555 important factor limiting NPF events at the kerbside site, SO<sub>2</sub> was found to have smaller  
556 concentrations on event days for all areas, which indicates that either on average it is in sufficient  
557 concentrations for NPF events to occur, or that other variables that participate in the production  
558 mechanism of H<sub>2</sub>SO<sub>4</sub> are more important. The growth rate of the newly formed particles increases  
559 from the rural site to the kerbside and is greater in summer compared to other seasons for all three

560 sites. Almost half of the NPF events at the rural and urban background sites were found to happen  
561 simultaneously. In these cases, the atmospheric conditions were cleaner, which resulted in slower  
562 growth rates. While most of the chemical species available were at lower concentrations in regional  
563 events, a difference in the behaviour with respect to sulphate and organic compounds was found  
564 between the two background site types.

565

566 The prevailing origin of air masses in the southern U.K. is from mid and high latitudes of the Atlantic  
567 Ocean. These fast-moving air masses present an increased probability for NPF to occur. The case is  
568 similar for the cooler and cleaner arctic air masses, while air masses from the tropics and continental  
569 Europe, having greater pollutant content, have decreased NPF probability, but a higher growth rate  
570 of particles when NPF events occurred. Regional events appear to be more associated with cleaner  
571 air masses, presenting a smaller growth rate and condensation sink compared to local events. The  
572 difference in growth rate is probably related to the greater content of condensable species; a positive  
573 relation of particle survival probability with temperature was also found.

574

575 Comparing the background areas in this study, particles of 16-20 nm were found to be about 20%  
576 greater in concentration (above long-term average) on NPF event days at the urban background site  
577 compared with the rural site. This is associated with a higher abundance of condensable species in  
578 the urban environment, which enhances the nucleation and growth process. This effect though is  
579 limited as particle size increases and NPF events have a greater effect on the overall  $N_{<100\text{ nm}}$  in the

580 rural areas, compared to urban, as calculated by the NSF. The effect becomes even smaller at the  
581 kerbside as the number of background particles emitted by traffic is a lot greater.

582  
583 The occurrence of NPF events at the highly polluted Marylebone Road site is at first sight surprising  
584 given the elevated condensation sink. This must be counteracted by an abundance of condensable  
585 material, which is surprising given the generally modest rate of atmospheric oxidation processes in  
586 comparison to residence times in a street canyon (Harrison, 2017). However, Giorio et al. (2015),  
587 using Aerosol Time-of-Flight Mass Spectrometry, reported rapid chemical processes within the  
588 Marylebone Road street canyon leading to production of secondary particulate matter from road  
589 traffic emissions. They postulated that this resulted from very local gas to particle conversion from  
590 vehicle-emitted pollutants. Condensation of such reaction products upon pre-existing particles could  
591 explain the enhanced particle growth rates observed at Marylebone Road (Figure 3).

592  
593 Finally, particle survival probability was found to decrease moving from rural to urban areas. While  
594 formation and initial growth of new particles is increased in urban areas, their survivability reduces  
595 as their size increases. The probability of particles to survive to greater sizes was found to be increased  
596 in summer for all areas, which is also explained by the higher growth rate. The probability is also  
597 different depending upon the origin of the air masses and is related to conditions specific for each  
598 area.

599

600 In the present work, the effects of atmospheric conditions upon the NPF process are studied. NPF is  
601 a complex process, highly affected by meteorological conditions (local and synoptic), the chemical  
602 composition as well as the pre-existing conditions in an area. For this reason, the study of NPF events  
603 in one area cannot provide safe assumptions for other areas, as the mixture of conditions found in  
604 different places is unique and alters the occurrence and development of NPF events. Thus, more  
605 studies on the conditions and the trends in NPF events should be conducted to better understand the  
606 effect of the numerous variables that affect those processes.

607

#### 608 **DATA AVAILABILITY**

609 Data supporting this publication are openly available from the UBIRA eData repository at  
610 <https://doi.org/10.25500/edata.bham.00000307>.

611

#### 612 **AUTHOR CONTRIBUTIONS**

613 This study was conceived by MD and RMH who also contributed to the final manuscript. The data  
614 analysis was carried out by DB with guidance from DCSB, and DB also prepared the first draft of  
615 the manuscript. FDP provided advice on the analysis.

#### 616 **COMPETING INTERESTS**

617 The authors have no conflict of interests.

618

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623

## 624 REFERENCES

- 625 Alam, A., Shi, J. P. and Harrison, R. M.: Observations of new particle formation in urban air, *Journal*  
626 *of Geophysical Research: Atmospheres*, 108, 4093–4107, doi:10.1029/2001JD001417, 2003.
- 627
- 628 Atkinson, R. W., Fuller, G. W., Anderson, H. R., Harrison, R. M. and Armstrong, B.: Urban ambient  
629 particle metrics and health: A time-series analysis, *Epidemiology*, 21, 501–511, 2010.
- 630
- 631 [Beccaceci, S., McGhee, E., Robins, C., Butterfield, D., Tompkins, J., Quincey, P., Brown, R., Green,](http://uk-air.defra.gov.uk/library/reports?section_id=13)  
632 [D., Tremper, A., Priestman, M., Font Font, A.: Airborne particulate concentrations and numbers in](http://uk-air.defra.gov.uk/library/reports?section_id=13)  
633 [the United Kingdom \(phase 3\): Annual report 2015, available at http://uk-](http://uk-air.defra.gov.uk/library/reports?section_id=13)  
634 [air.defra.gov.uk/library/reports?section\\_id=13](http://uk-air.defra.gov.uk/library/reports?section_id=13)
- 635
- 636 Beddows, D. C. S., Harrison, R. M., Green, D. C., and Fuller, G. W.: Receptor modelling of both  
637 particle composition and size distribution from a background site in London, UK, *Atmos. Chem.*  
638 *Phys.*, 15, 10107–10125, 2015.
- 639
- 640 Berndt, T., Böge, O., and Stratmann, F.: Formation of atmospheric H<sub>2</sub>SO<sub>4</sub>H<sub>2</sub>O particles in the  
641 absence of organics: A laboratory study, *Geophys. Res. Lett.*, 33, 2–6, 2006.
- 642
- 643 Bianchi, F., Trostl, J., Junninen, H., Frege, C., Henne, S., Hoyle, C. R., Molteni, U., Herrmann, E.,  
644 Adamov, A., Bukowiecki, N., Chen, X., Duplissy, J., Gysel, M., Hutterli, M., Kangasluoma, J.,  
645 Kontkanen, J., Kurten, A., Manninen, H. E., Munch, S., Perakyla, O., Petaja, T., Rondo, L.,  
646 Williamson, C., Weingartner, E., Curtius, J., Worsnop, D. R., Kulmala, M., Dommen, J., and  
647 Baltensperger, U.: New particle formation in the free troposphere: A question of chemistry and  
648 timing, *Science*, 352, 1109–1112, 2016.
- 649
- 650 Bigi, A. and Harrison, R. M.: Analysis of the air pollution climate at a central urban background site,  
651 *Atmos. Environ.*, 44, 2004–2012, 2010.
- 652
- 653 Brines, M., Dall'Osto, M., Beddows, D. C. S., Harrison, R. M., and Querol, X.: Simplifying aerosol  
654 size distributions modes simultaneously detected at four monitoring sites during SAPUSS, *Atmos.*  
655 *Chem. Phys.*, 14, 2973–2986, 2014.
- 656
- 657 Carslaw, D. C. and Ropkins, K.: openair — An R package for air quality data analysis, *Environ.*  
658 *Modell. Softw.*, 27–28, 52–61, 2012.
- 659
- 660 Charron, A., Birmili, W., and Harrison, R. M.: Factors influencing new particle formation at the rural  
661 site, Harwell, United Kingdom, *J. Geophys. Res., Atmospheres*, 112,  
662 doi:10.1029/2007JD0084252007.

663  
664 Charron, A., Birmili, W., and Harrison, R. M.: Fingerprinting particle origins according to their size  
665 distribution at a UK rural site, *J. Geophys. Res., Atmospheres*, 113, D07202,  
666 doi:10.1029/2007JD008562, 2008.

667 Charron, A., Degrendele, C., Laongsri, B., and Harrison, R. M.: Receptor modelling of secondary  
668 and carbonaceous particulate matter at a southern UK site, *Atmos. Chem. Phys.*, 13, 1879–1894,  
669 2013.

670  
671 Charron, A. and Harrison, R. M.: Primary particle formation from vehicle emissions during exhaust  
672 dilution in the roadside atmosphere, *Atmos. Environ.*, 37, 4109–4119, 2003.

673  
674 Dal Maso, M., Kulmala, M., Lehtinen, K. E. J., Mäkelä, J. M., Aalto, P., and O’Dowd, C. D.:  
675 Condensation and coagulation sinks and formation of nucleation mode particles in coastal and boreal  
676 forest boundary layers, *J. Geophys. Res., Atmospheres*, 107, doi: 10.1029/2001JD001053, 2002.

677  
678 Dal Maso, M., Kulmala, M., Riipinen, I., Wagner, R., Hussein, T., Aalto, P. P., and Lehtinen, K. E.  
679 J.: Formation and growth of fresh atmospheric aerosols: Eight years of aerosol size distribution data  
680 from SMEAR II, Hyytiälä, Finland, *Boreal Environ. Res.*, 10, 323–336, 2005.

681  
682 Dall’Osto, M., Beddows, D. C. S., Asmi, A., Poulain, L., Hao, L., Freney, E., Allan, J. D.,  
683 Canagaratna, M., Crippa, M., Bianchi, F., de Leeuw, G., Eriksson, A., Swietlicki, E., Hansson, H. C.,  
684 Henzing, J. S., Granier, C., Zemankova, K., Laj, P., Onasch, T., Prevot, A., Putaud, J. P., Sellegri, K.,  
685 Vidal, M., Virtanen, A., Simo, R., Worsnop, D., O’Dowd, C., Kulmala, M., and Harrison, R. M.:  
686 Novel insights on new particle formation derived from a pan-european observing system, *Sci. Rep.*,  
687 8, 1482, 2018.

688  
689 Dall’Osto, M., Beddows, D. C. S., Pey, J., Rodriguez, S., Alastuey, A., M. Harrison, R., and Querol,  
690 X.: Urban aerosol size distributions over the Mediterranean city of Barcelona, NE Spain, *Atmos.*  
691 *Chem. Phys.*, 12, 10693–10707, 2012.

692  
693 Dall’Osto, M., Querol, X., Alastuey, A., O’Dowd, C., Harrison, R. M., Wenger, J., and Gómez-  
694 Moreno, F. J.: On the spatial distribution and evolution of ultrafine particles in Barcelona, *Atmos.*  
695 *Chem. Phys.*, 13, 741–759, 2013.

696  
697 Dall’Osto, M., Thorpe, A., Beddows, D. C. S., Harrison, R. M., Barlow, J. F., Dunbar, T., Williams,  
698 P. I., and Coe, H.: Remarkable dynamics of nanoparticles in the urban atmosphere, *Atmos. Chem.*  
699 *Phys.*, 11, 6623–6637, 2011.

700  
701 Dall’Osto, M., Beddows, D. C. S., Tunved, P., Krejci, R., Ström, J., Hansson, H. C., Yoon, Y. J.,  
702 Park, K. T., Becagli, S., Udisti, R., Onasch, T., Ódowd, C. D., Simó, R., and Harrison, R. M.: Arctic

703 sea ice melt leads to atmospheric new particle formation, *Sci.Rep.*, 7, 0–10, 2017.  
704  
705  
706 Dameto de España, C., Wonaschütz, A., Steiner, G., Rosati, B., Demattio, A., Schuh, H. and  
707 Hitzengerger, R.: Long-term quantitative field study of New Particle Formation (NPF) events as a  
708 source of Cloud Condensation Nuclei (CCN) in the urban background of Vienna, *Atmos. Environ.*,  
709 164, 289–298, 2017.  
710  
711 Davidson, C. I., Phalen, R. F., and Solomon, P. A.: Airborne particulate matter and human health: A  
712 review, *Aerosol Sci, Technol.*, 39, 737–749, 2005.  
713  
714 Draxler, R. R., and Hess, G. D.: An Overview of the HYSPLIT\_4 Modelling System for Trajectories,  
715 Dispersion, and Deposition, *Australian Meteorolog. Mag.*, 47, 295–308, 1998.  
716  
717 Ehn, M., Vuollekoski, H., Petäjä, T., Kerminen, V.-M., Vana, M., Aalto, P., de Leeuw, G., Ceburnis,  
718 D., Dupuy, R., O’Dowd, C. D., and Kulmala, M.: Growth rates during coastal and marine new particle  
719 formation in western Ireland, *J. Geophys. Res.*, 115, D18218,  
720 <http://dx.doi.org/10.1029/2010JD014292>, 2010.  
721  
722 Fiedler, V., Dal Maso, M., Boy, M., Aufmhoff, H., Hoffmann, J., Schuck, T., Birmili, W., Arnold,  
723 F., and Kulmala, M.: The contribution of sulphuric acid to atmospheric particle formation and growth:  
724 a comparison between boundary layers in Northern and Central Europe, *Atmos. Chem. Phys. Discuss.*  
725 5, 573–605, 2005.  
726  
727 Fuchs, N. A. and Sutugin, A. G.: Highly Dispersed Aerosols, *Foreign Sci. and Technol. Center*, 1–  
728 86, 1971.  
729  
730 Gentner, D. R., Isaacman, G., Worton, D. R., Chan, A. W. H., Dallmann, T. R., Davis, L., Liu, S.,  
731 Day, D. A., Russell, L. M., Wilson, K. R., Weber, R., Guha, A., Harley, R. A., and Goldstein, A. H.:  
732 Elucidating secondary organic aerosol from diesel and gasoline vehicles through detailed  
733 characterization of organic carbon emissions, *Proc. Natl. Acad. Sci.*, 109, 18318–18323, 2012.  
734  
735 Hama, S. M. L., Cordell, R. L., Kos, G. P. A., Weijers, E. P., and Monks, P. S.: Sub-micron particle  
736 number size distribution characteristics at two urban locations in Leicester, *Atmos. Res.*, 194, 1–16,  
737 2017.  
738  
739 Harrison, R. M., Beddows, D. C. S., Alam, M. S., Singh, A., Brean, J., and R. Xu: Interpretation of  
740 Particle number size distributions measured across an urban area during the FASTER campaign, in  
741 preparation, 2018.  
742

743 Harrison, R. M.: Urban atmospheric chemistry: a very special case for study, *npj Climate and Atmos.*  
 744 *Sci.*, 1, 5, 2017.  
 745  
 746 Harrison, R. M., Shi, J. P., Xi, S., Khan, A., Mark, D., Kinnersley, R., and Yin, J.: Measurement of  
 747 number, mass and size distribution of particles in the atmosphere, *Philos. Trans. A. Math. Phys. Eng.*  
 748 *Sci.*, 358, 2567–2580, 2000.  
 749  
 750 Harrison, R. M. and Yin, J.: Particulate matter in the atmosphere: Which particle properties are  
 751 important for its effects on health?, *Sci. Tot. Environ.*, 249, 85–101, 2000.  
 752  
 753 Harrison, R.M., Yin, J., Tilling, R.M., Cai, X., Seakins, P.W., Hopkins, J.R., Lansley, D.L.,  
 754 Lewis, A.C., Hunter, M.C., Heard, D.E., Carpenter, L.J., Creasey, D.C., Lee, J.D., Pilling, M.J.,  
 755 Carslaw, N., Emmerson, K.M., Redington, A., Derwent, R.G., Ryall, D., Mills G., and Penkett, S.A.,  
 756 Measurement and Modelling of Air Pollution and Atmospheric Chemistry in the UK West  
 757 Midlands Conurbation: Overview of the PUMA Consortium Project, *Sci. Tot. Environ.*, 360, 5-25  
 758 2006.  
 759  
 760 Hietikko, R., Kuuluvainen, H., Harrison, R. M., Portin, H., Timonen, H., Niemi, J. V., Ronkko, T.:  
 761 Diurnal variation of nanocluster aerosol concentrations and emission factors in a street canyon,  
 762 *Atmos. Environ.*, 189, 98-106, 2018.  
 763  
 764 Iida, K., Stolzenburg, M. R., McMurry, P. H., and Smith, J. N.: Estimating nanoparticle growth rates  
 765 from size-dependent charged fractions: Analysis of new particle formation events in Mexico City, *J.*  
 766 *Geophys. Res. Atmospheres*, 113, D05207, doi:10.1029/2007JD009260, 2008.  
 767  
 768 Jayaratne, R., Pushpawela, B., He, C., Li, H., Gao, J., Chai, F., and Morawska, L.: Observations of  
 769 particles at their formation sizes in Beijing, China, *Atmos. Chem. Phys.*, 17, 8825–8835, 2017.  
 770  
 771 Jeong, C.-H., Evans, G. J., McGuire, M. L., Chang, R. Y.-W., Abbatt, J. P. D., Zeromskiene, K.,  
 772 Mozurkewich, M., Li, S.-M., and Leaitch, W. R.: Particle formation and growth at five rural and  
 773 urban sites, *Atmos. Chem. Phys.*, 10, 7979–7995, 2010.  
 774  
 775 Kalkavouras, P., Bossioli, E., Bezantakos, S., Bougiatioti, A., Kalivitis, N., Stavroulas, I.,  
 776 Kouvarakis, G., Protonotariou, A. P., Dandou, A., Biskos, G., Mihalopoulos, N., Nenes, A., and  
 777 Tombrou, M.: New particle formation in the southern Aegean Sea during the Etesians: Importance  
 778 for CCN production and cloud droplet number, *Atmos. Chem. Phys.*, 17, 175–192, 2017.  
 779  
 780 Kecorius, S., Kivekäs, N., Kristensson, A., Tuch, T., Covert, D. S., Birmili, W., Lihavainen, H.,  
 781 Hyvärinen, A. P., Martinsson, J., Sporre, M. K., Swietlicki, E., Wiedensohler, A., and Ulevicius, V.:  
 782 Significant increase of aerosol number concentrations in air masses crossing a densely trafficked sea

783 area, *Oceanologia*, 58, 1–12, 2016.

784

785

786 Kelly, F. J. and Fussell, J. C.: Size, source and chemical composition as determinants of toxicity

787 attributable to ambient particulate matter, *Atmos. Environ.*, 60, 504–526, 2012.

788

789 Keuken, M. P., Moerman, M., Zandveld, P., Henzing, J. S., and Hoek, G.: Total and size-resolved

790 particle number and black carbon concentrations in urban areas near Schiphol airport (the

791 Netherlands), *Atmos. Environ.*, 104, 132–142, 2015.

792

793 Kirkby, J. et al.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol

794 nucleation, *Nature*, 476(7361), pp. 429–435, 2011.

795

796 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagne, S., Ickes,

797 L., Kurten, A., Kupc, Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G.,

798 Wimmer, D., Amorim, A. A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Downard, A.,

799 Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud, W., Junninen, H., Kreiss, F., Kvashin,

800 A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R., Makhmutov, V., Mathot, S., Mikkila, J.,

801 Minginette, P., Mogo, S., Nieminen, T., Onnela, A., Pereira, P., Petaja, T., Schnitzhofer, R., Seinfeld,

802 J. H., Sipila, M., Stozhkov, Y., Stratmann, F., Tome, A., Vanhanen, J., Viisanen, Y., Vrtala, A.,

803 Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D.

804 R., Baltensperger, U., and Kulmala M.: Ion-induced nucleation of pure biogenic particles, *Nature*. Nature Publishing Group, 533, 521–526,

805 2016.

806

807 Korhonen, P., Kulmala, M., Laaksonen, A., Viisanen, Y., Mcgraw, R., and Seinfeld, J. H.: Ternary

808 nucleation of  $\text{H}_2\text{SO}_4$ ,  $\text{NH}_3$  and  $\text{H}_2\text{O}$  in the atmosphere, *J. Geophys. Res.*, 104, D21, 26,349–26,353,

809 1999.

810

811 Kuang, C., McMurry, P. H., McCormick, A. V., and Eisele, F. L.: Dependence of nucleation rates on

812 sulfuric acid vapor concentration in diverse atmospheric locations, *J. Geophys. Res.*, Atmospheres,

813 113, D10209, doi:10.1029/2007JD009253, 2008.

814

815 Kulmala, M., Dal Maso, M., Mäkelä, J. M., Pirjola, L., Väkevä, M., Aalto, P., Mäkelä, P., Hämeri, K., and O'Dowd, C. D.: On the formation, growth and composition of nucleation mode

816 particles, *Tellus, Series B: Chem. Phys.Meteorol.*, 53, 479–490, 2001.

817

818 Kulmala, M., Kerminen, V.-M., Petäjä, T., Ding, A. J., and Wang, L.: Atmospheric gas-to-particle

819 conversion: why NPF events are observed in megacities?, *Faraday Discuss.*, 271–288, 2017.

820

821

822

823 Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petaja, T.,  
 824 Sipila, M., Schobesberger, S., Rantala, P., Franchin, A., Jokinen, T., Jarvinen, E., Aijala, M.,  
 825 Kangasluoma, J., Hakala, J., Aalto, P. P., Paasonen, P., Mikkila, J., Vanhanen, J., Aalto, J., Hakola,  
 826 H., Makkonen, U., Ruuskanen, T., Mauldin, R. L., Duplissy, J., Vehkamäki, H., Back, J., Kortelainen,  
 827 A., Riipinen, I., Kurten, T., Johnston, M. V., Smith, J. N., Ehn, M., Mentel, T. F., Lehtinen, K. E. J.,  
 828 Laaksonen, A., Kerminen, V.-M., and Worsnop, D. R.: Direct Observations of Atmospheric Aerosol  
 829 Nucleation, *Science*, 339, 943–946, 2013.  
 830  
 831 Kulmala, M., Luoma, K., Virkkula, A., Petäjä, T., Paasonen, P., Kerminen, V. M., Nie, W., Qi, X.,  
 832 Shen, Y., Chi, X., and Ding, A.: On the mode-segregated aerosol particle number concentration load:  
 833 Contributions of primary and secondary particles in Hyytiälä and Nanjing, *Boreal Environ. Res.*, 21,  
 834 319–331, 2016.  
 835  
 836 Kulmala, M., Petäjä, T., Ehn, M., Thornton, J., Sipilä, M., Worsnop, D. R., and Kerminen, V.-M.:  
 837 Chemistry of Atmospheric Nucleation: On the Recent Advances on Precursor Characterization and  
 838 Atmospheric Cluster Composition in Connection with Atmospheric New Particle Formation,  
 839 *Ann.Rev.Phys. Chem.*, 65, 21–37, 2014.  
 840  
 841 Kulmala, M., Petäjä, T., Mönkkönen, P., Koponen, I. K., Dal Maso, M., Aalto, P. P., Lehtinen, K. E.  
 842 J., and Kerminen, V.-M.: On the growth of nucleation mode particles: source rates of condensable  
 843 vapor in polluted and clean environments, *Atmos. Chem. Phys. Discuss.*, 4, 6943–6966, 2005.  
 844  
 845 Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M.,  
 846 Aalto, P. P., Junninen, H., Paasonen, P., Riipinen, I., Lehtinen, K. E. J., Laaksonen, A., and Kerminen,  
 847 V. M.: Measurement of the nucleation of atmospheric aerosol particles, *Nature Protocols*, 7, 1651–  
 848 1667, 2012.  
 849  
 850 Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili, W., and  
 851 McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: A review of  
 852 observations, *J. Aerosol Sci.*, 35, 143–176, 2004.  
 853  
 854 Laaksonen, A., Kulmala, M., O’Dowd, C. D., Joutsensaari, J., Vaattovaara, P., Mikkonen, S.,  
 855 Lehtinen, K. E. J., Sogacheva, L., Dal Maso, M., Aalto, P., Petäjä, T., Sogachev, A., Yoon, Y. J.,  
 856 Lihavainen, H., Nilsson, D., Facchini, M. C., Cavalli, F., Fuzzi, S., Hoffmann, T., Arnold, F., Hanke,  
 857 M., Sellegri, K., Umann, B., Junkermann, W., Coe, H., Allan, J. D., Alfarra, M. R., Worsnop, D. R.,  
 858 Riekkola, M. L., Hyötyläinen, T., and Viisanen, Y.: The role of VOC oxidation products in  
 859 continental new particle formation, *Atmos. Chem. Phys.*, 8, 657–2665, 2008.  
 860  
 861 Ma, N. and Birmili, W.: Estimating the contribution of photochemical particle formation to ultrafine  
 862 particle number averages in an urban atmosphere, *Sci. Tot. Environ.*, 512–513, 154–166, 2015.

863  
864  
865 MacNee, W. and Donaldson, K.: Mechanism of lung injury caused by PM10 and ultrafine particles  
866 with special reference to COPD, *Europ. Respirat. J.*, 21, 47S–51S, 2003.  
867  
868 Makkonen, R., Asmi, A., Kerminen, V. M., Boy, M., Arneth, A., Hari, P., and Kulmala, M.: Air  
869 pollution control and decreasing new particle formation lead to strong climate warming, *Atmos.*  
870 *Chem. Phys.*, 12, 1515–1524, 2012.  
871  
872 Masiol, M., Harrison, R. M., Vu, T. V., and Beddows, D. C. S.: Sources of sub-micrometre particles  
873 near a major international airport, *Atmos. Chem. Phys.*, 17, 12379–12403, 2017.  
874  
875 Merikanto, J., Spracklen, D. V., Mann, G. W., Pickering, S. J., and Carslaw, K. S.: Impact of  
876 nucleation on global CCN, *Atmos.Chem. Phys.*, 9, 8601–8616, 2009.  
877  
878 Metzger, A., Verheggen, B., Dommen, J., Duplissy, J., Prevot, A. S. H., Weingartner, E., Riipinen,  
879 I., Kulmala, M., Spracklen, D. V., Carslaw, K. S., and Baltensperger, U.: Evidence for the role of  
880 organics in aerosol particle formation under atmospheric conditions, *Proc. Natl. Acad. Sci.*, 107,  
881 6646–6651, 2010.  
882  
883 Minguillón, M. C., Brines, M., Pérez, N., Reche, C., Pandolfi, M., Fonseca, A. S., Amato, F.,  
884 Alastuey, A., Llyasota, A., Codina, B., Lee, H. K., Eun, H. R., Ahn, K. H., and Querol, X.: New  
885 particle formation at ground level and in the vertical column over the Barcelona area, *Atmos. Res.*,  
886 164–165, 118–130, 2015.  
887  
888 Napari, I., Noppel, M., Vehkamäki, H., and Kulmala, M.: An improved model for ternary nucleation  
889 of sulfuric acid-ammonia-water, *J. Chem. Phys.*, 116, 4221–4227, 2002.  
890  
891 Németh, Z., Rosati, B., Zíková, N., Salma, I., Bozó, L., Dameto de España, C., Schwarz, J., Ždímal,  
892 V., and Wonaschütz, A.: Comparison of atmospheric new particle formation events in three Central  
893 European cities, *Atmos. Environ.*, 178, 191–197, 2018.  
894  
895 Németh, Z. and Salma, I.: Spatial extension of nucleating air masses in the Carpathian Basin, *Atmos.*  
896 *Chem. Phys.*, 14, 8841–8848, 2014.  
897  
898 Nieminen, T., Kerminen, V.-M., Petäjä, T., Aalto, P. P., Arshinov, M., Asmi, E., Baltensperger, U.,  
899 Beddows, D. C. S., Beukes, J. P., Collins, D., Ding, A., Harrison, R. M., Henzing, B., Hooda, R., Hu,  
900 M., Hörrak, U., Kivekäs, N., Komsaare, K., Krejci, R., Kristensson, A., Laakso, L., Laaksonen, A.,  
901 Leaitch, W. R., Lihavainen, H., Mihalopoulos, N., Németh, Z., Nie, W., O ’dowd, C., Salma, I.,  
902 Sellegri, K., Svenningsson, B., Swietlicki, E., Tunved, P., Ulevicius, V., Vakkari, V., Vana, M.,

Wiedensohler, A., Wu, Z., Virtanen, A., and Kulmala, M.: Global analysis of continental boundary layer new particle formation based on long-term measurements, *Atmos. Chem. Phys. Discuss.*, 5194, 2018–304, 2018.

Nilsson, E. D., Paatero, J. and Boy, M.: Effects of air masses and synoptic weather on aerosol formation in the continental boundary layer, *Tellus, Series B: Chem. Phys. Meteorol.*, 53, 462–478, 2001.

O’Dowd, C. D., Aalto, P., Hmeri, K., Kulmala, M., and Hoffmann, T.: Atmospheric particels from organic vapours, *Nature*, 416, 497–498, 2002.

O’Dowd, C., Jimenez, J. L., Bahreini, R., Flagan, R. C., Seinfeld, J. H., Hameri Kaarle, Pirjola, L., Kulmala, M., Gerard Jennings, S., and Hoffmann, T.: Marine aerosol formation from biogenic iodine emissions, *Nature*, 417, 1–5, 2002.

Oberdurst, G.: Toxicology of ultrafine particles: in vivo studies, *Philos. Trans. A. Math. Phys. Eng. Sci.*, 358, 2719–2740, 2000.

Park, M., Yum, S. S., and Kim, J. H.: Characteristics of submicron aerosol number size distribution and new particle formation events measured in Seoul, Korea, during 2004–2012, *Asia-Pacific J. Atmos. Sci.*, 51, 1–10, 2015.

Peng, Y., Dong, Y., Li, X., Liu, X., Dai, J., Chen, C., Dong, Z., Du, C., and Wang, Z.: Different Characteristics of New Particle Formation Events at Two Suburban Sites in Northern China, *Atmosphere*, 8, 58, 2017.

Penttinen, P., Timonen, K. L., Tiittanen, P., Mirme, A., Ruuskanen, J., and Pekkanen, J.: Number concentration and size of particles in urban air: Effects on spirometric lung function in adult asthmatic subjects, *Environ. Health Perspect.*, 109, 319–323, 2001.

Petäjä, T., Mauldin, R. L., III, Kosciuch, E., McGrath, J., Nieminen, T., Paasonen, P., Boy, M., Adamov, A., Kotiaho, T., and Kulmala, M.: Sulfuric acid and OH concentrations in a boreal forest site, *Atmos. Chem. Phys.*, 9, 7435–7448, 2009.

Pikridas, M., Sciare, J., Freutel, F., Crumeyrolle, S., Von Der Weiden-Reinmüller, S. L., Borbon, A., Schwarzenboeck, A., Merkel, M., Crippa, M., Kostenidou, E., Psichoudaki, M., Hildebrandt, L., Engelhart, G. J., Petäjä, T., Prévôt, A. S. H., Drewnick, F., Baltensperger, U., Wiedensohler, A., Kulmala, M., Beekmann, M., and Pandis, S. N.: In situ formation and spatial variability of particle number concentration in a European megacity, *Atmos. Chem. Phys.*, 15, 0219–10237, 2015.

943 Politis, M., Pilinis, C., and Lekkas, T. D.: Ultrafine particles (UFP) and health effects. Dangerous.  
 944 Like no other PM? Review and analysis, *Global Nest J.*, 10, 439–452, 2008.

945 Rahman, M. M., Mazaheri, M., Clifford, S., and Morawska, L.: Estimate of main local sources to  
 946 ambient ultrafine particle number concentrations in an urban area, *Atmos. Res.*, 194, 178–189, 2017.  
 947

948 Riccobono, F., Schobesberger, S., Scott, C. E., Dommen, J., Ortega, I. K., Rondo, L., Almeida, J.,  
 949 Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Downard, A., Dunne, E. M., Duplissy, J.,  
 950 Ehrhart, S., Flagan, R. C., Franchin, A., Hansel, A., Junninen, H., Kajos, M., Keskinen, H., Kupc, A.,  
 951 Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Petäjä, T., Tsagkogeorgas, G., Vaattovaara,  
 952 P., Viisanen, Y., Vrtala, A., and Wagner, P. E.: Oxidation Products of Biogenic Atmospheric  
 953 Particles, *Science*, 717, 17–722, 2014.  
 954

955 Riipinen, I., Sihto, S.-L., Kulmala, M., Arnold, F., Dal Maso, M., Birmili, W., Saarnio, K., Teinilä,  
 956 K., Kerminen, V.-M., Laaksonen, A., and Lehtinen, K. E. J.: Connections between atmospheric  
 957 sulphuric acid and new particle formation during QUEST III–IV campaigns in Heidelberg and  
 958 Hyytiälä, *Atmos. Chem. Phys.*, 7, 1899–1914, 2007.  
 959

960 Robinson, A. L., Donahue, N. M., Shrivastava, M. K., Weitkamp, E. A., Sage, A. M., Grieshop, A.  
 961 P., Lane, T. E., Pierce, J. R., and Pandis, S. N.: Rethinking Organic Aerosols :, *Science*, 315, 1259–  
 962 1262, 2007.  
 963

964 Rönkkö, T., Kuuluvainen, H., Karjalainen, P., Keskinen, J., Hillamo, R., Niemi, J. V., Pirjola, L.,  
 965 Timonen, H. J., Saarikoski, S., Saukko, E., Järvinen, A., Silvennoinen, H., Rostedt, A., Olin, M., Yli-  
 966 Ojanperä, J., Nousiainen, P., Kousa, A. and Dal Maso, M.: Traffic is a major source of atmospheric  
 967 nanocluster aerosol, *Proc. Natl. Acad. Sci.*, 114, 7549–7554, 2017.  
 968

969 Salma, I., Borsós, T., Németh, Z., Weidinger, T., Aalto, P., and Kulmala, M.: Comparative study of  
 970 ultrafine atmospheric aerosol within a city, *Atmos. Environ.*, 92, 154–161, 2014.  
 971

972 Salma, I., Németh, Z., Kerminen, V. M., Aalto, P., Nieminen, T., Weidinger, T., Molnár, Á., Imre,  
 973 K., and Kulmala, M.: Regional effect on urban atmospheric nucleation, *Atmos. Chem. Phys.*, 16,  
 974 8715–8728, 2016.  
 975

976 Salma, I., Varga, V., and Németh, Z.: Quantification of an atmospheric nucleation and growth process  
 977 as a single source of aerosol particles in a city, *Atmos. Chem. Phys.*, 17, 15007–15017, 2017.  
 978

979 Samoli, E., Atkinson, R. W., Analitis, A., Fuller, G. W., Beddows, D., Green, D. C., Mudway, I. S.,  
 980 Harrison, R. M., Anderson, H. R., and Kelly, F. J.: Differential health effects of short-term exposure  
 981 to source-specific particles in London, U.K., *Environ. Intl.*, 97, 246–253, 2016.  
 982

Seinfeld, J. H. and Pandis, S. N.: Atmospheric Chemistry and Physics: From Air Pollution to Climate Change, 3rd Ed. New Jersey, Canada, John Wiley & Sons, Inc, 2012.

Shen, X., Sun, J., Kivekäs, N., Kristensson, A., Zhang, X., Zhang, Y., Zhang, L., Fan, R., Qi, X., Ma, Q. and Zhou, H.: Spatial distribution and occurrence probability of regional new particle formation events in eastern China, *Atmos. Chem. Phys.*, 18, 587–599, 2018.

Shi, J. P., Evans, D. E., Khan, A. A., and Harrison, R. M.: Sources and concentration of nanoparticles (<10nm diameter) in the urban atmosphere, *Atmos. Environ.*, 35, 1193–1202, 2001.

Shi, J. P. and Harrison, R. M.: Investigation of ultrafine particle formation during diesel exhaust dilution, *Environ. Sci. Technol.*, 33, 3730–3736, 1999.

Spracklen, D. V., Carslaw, K. S., Kulmala, M., Kerminen, V. M., Sihto, S. L., Riipinen, I., Merikanto, J., Mann, G. W., Chipperfield, M. P., Wiedensohler, A., Birmili, W., and Lihavainen, H.: Contribution of particle formation to global cloud condensation nuclei concentrations, *Geophys. Res. Lett.*, 35, 1–5, 2008.

Spracklen, D. V., Carslaw, K. S., Merikanto, J., Mann, G. W., Reddington, C. L., Pickering, S., Ogren, J. A., Andrews, E., Baltensperger, U., Weingartner, E., Boy, M., Kulmala, M., Laakso, L., Lihavainen, H., Kivekäs, N., Komppula, M., Mihalopoulos, N., Kouvarakis, G., Jennings, S. G., O'Dowd, C., Birmili, W., Wiedensohler, A., Weller, R., Gras, J., Laj, P., Sellegri, K., Bonn, B., Krejci, R., Laaksonen, A., Hamed, A., Minikin, A., Harrison, R. M., Talbot, R., and Sun, J.: Explaining global surface aerosol number concentrations in terms of primary emissions and particle formation, *Atmos. Chem. Phys.*, 10, 4775–4793, 2010.

Sutton, M. A., Place, C. J., Eager, M., Fowler, D., and Smith, R. I.: Assessment of the magnitude of ammonia emissions in the UK, *Atmos. Environ.*, 29, 1393–1411, 1995.

Tröstl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege, C., Bianchi, F., Wagner, R., Simon, M., Lehtipalo, K., Williamson, C., Craven, J. S., Duplissy, J., Adamov, A., Almeida, J., Bernhammer, A.-K., Breitenlechner, M., Brilke, S., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Gysel, M., Hansel, A., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Keskinen, H., Kim, J., Krapf, M., Kürten, A., Laaksonen, A., Lawler, M., Leiminger, M., Mathot, S., Möhler, O., Nieminen, T., Onnela, A., Petäjä, T., Piel, F. M., Miettinen, P., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Sengupta, K., Sipilä, M., Smith, J. N., Steiner, G., Tomè, A., Virtanen, A., Wagner, A. C., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Carslaw, K. S., Curtius, J., Dommen, J., Kirkby, J., Kulmala, M., Riipinen, I., Worsnop, D. R., Donahue, N. M., and Baltensperger, U.: The role of low-volatility organic compounds in initial particle growth in the atmosphere, *Nature*, 533, 527–531, 2016.

1023 Von Bismarck-Osten, C., Birmili, W., Ketzel, M., Massling, A., Petäjä, T., and Weber, S.:  
 1024 Characterization of parameters influencing the spatio-temporal variability of urban particle number  
 1025 size distributions in four European cities, *Atmos. Environ.*, 77, 415–429, 2013.  
 1026  
 1027 Wang, Z., Wu, Z., Yue, D., Shang, D., Guo, S., Sun, J., Ding, A., Wang, L., Jiang, J., Guo, H., Gao,  
 1028 J., Cheung, H. C., Morawska, L., Keywood, M., and Hu, M.: New particle formation in China: Current  
 1029 knowledge and further directions, *Sci. Tot. Environ.*, 258–266, 2017.  
 1030  
 1031 Wehner, B., Siebert, H., Stratmann, F., Tuch, T., Wiedensohler, A., Petäjä, T., Dal Maso, M., and  
 1032 Kulmala, M.: Horizontal homogeneity and vertical extent of new particle formation events, *Tellus,*  
 1033 *Series B: Chem. Phys. Meteorol.*, 59, 362–371, 2007.  
 1034  
 1035 Woo, K. S., Chen, D. R., Pui, D. Y. H., and McMurry, P. H.: Measurement of Atlanta aerosol size  
 1036 distributions: Observations of lutrafine particle events, *Aerosol Sci. Technol.*, 34, 5–87, 2001.  
 1037  
 1038 Xiao, S., Wang, M. Y., Yao, L., Kulmala, M., Zhou, B., Yang, X., Chen, J. M., Wang, D. F., Fu, Q.  
 1039 Y., Worsnop, D. R., and Wang, L.: Strong atmospheric new particle formation in winter in urban  
 1040 Shanghai, China, *Atmos. Chem. Phys.*, 15, 1769–1781, 2015.  
 1041  
 1042 Yli-Juuti, T., Nieminen, T., Hirsikko, A., Aalto, P. P., Asmi, E., Hörrak, U., Manninen, H. E.,  
 1043 Patokoski, J., Dal Maso, M., Petäjä, T., Rinne, J., Kulmala, M., and Riipinen, I.: Growth rates of  
 1044 nucleation mode particles in Hyytiälä during 2003–2009: Variation with particle size, season, data  
 1045 analysis method and ambient conditions, *Atmos. Chem. Phys.*, 11, 12865–12886, 2011.  
 1046  
 1047 Yue, D. L., Hu, M., Zhang, R. Y., Wang, Z. B., Zheng, J., Wu, Z. J., Wiedensohler, A., He, L. Y.,  
 1048 Huang, X. F., and Zhu, T.: The roles of sulfuric acid in new particle formation and growth in the  
 1049 mega-city of Beijing, *Atmos. Chem. Phys.*, 10, 4953–4960, 2010.  
 1050  
 1051 Zhang, X., Zhang, Y., Sun, J., Zheng, X., Li, G., and Deng, Z.: Characterization of particle number  
 1052 size distribution and new particle formation in an urban environment in Lanzhou, China, *J. Aerosol*  
 1053 *Sci.*, 103, 53–66, 2017.  
 1054  
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1061 **TABLE LEGENDS:**  
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1063 **Table 1:** Number of NPF events per site (in parenthesis the number of days with available  
1064 data).  
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1066 **Table 2:** Annual and seasonal NSF for all areas of study.  
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1070 **FIGURE LEGENDS:**  
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1072 **Figure 1:** Map of the measuring stations.  
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1074 **Figure 2:** Number of NPF events per season for all seven years of the present study (Winter –  
1075 DJF; Spring – MAM; Summer – JJA; Autumn – SON) at Harwell (rural), N.  
1076 Kensington (urban background) and Marylebone Road (urban roadside).  
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1078 **Figure 3:** Growth rate per season at the three sites.  
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1080 **Figure 4:** Diurnal variation of N<sub>16-20nm</sub> at each site: annual average-mean and NPF event days.  
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1082 **Figure 5:** Map and frequency of incoming air mass origin – average and for NPF events per site.  
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1084 **Figure 6:** Growth rate per incoming air mass at each of the sites.  
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1086 **Figure 7:** Survival parameter P (a) per season, (b) for regional and local events (for Marylebone  
1087 Road) is regional for all 3 sites and (c) by incoming air mass origin.  
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1089 **Table 1:** Number of NPF events per site [\(in parenthesis the number of days with available data\)](#).  
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	Harwell	N. Kensington	Marylebone Road	Regional (Background sites)*	Regional (All 3 sites)**
2009	9 (210)	0 (332)	4 (290)	0	0
2010	29 (262)	22 (310)	22 (292)	11	9
2011	15 (291)	10 (300)	23 (284)	4	1
2012	8 (334)	28 (303)	12 (140)	3	0
2013	25 (328)	23 (342)	27 <a href="#">(334)</a>	13	11
2014	29 (324)	34 (330)	13 <a href="#">(314)</a>	18	6
2015	25 (282)	22 (314)	18 <a href="#">(338)</a>	11	10
Overall	140 (2031)	139 (2231)	119 (1993)	60	37

1091 \* Refers to events occurring simultaneously at Harwell and N. Kensington

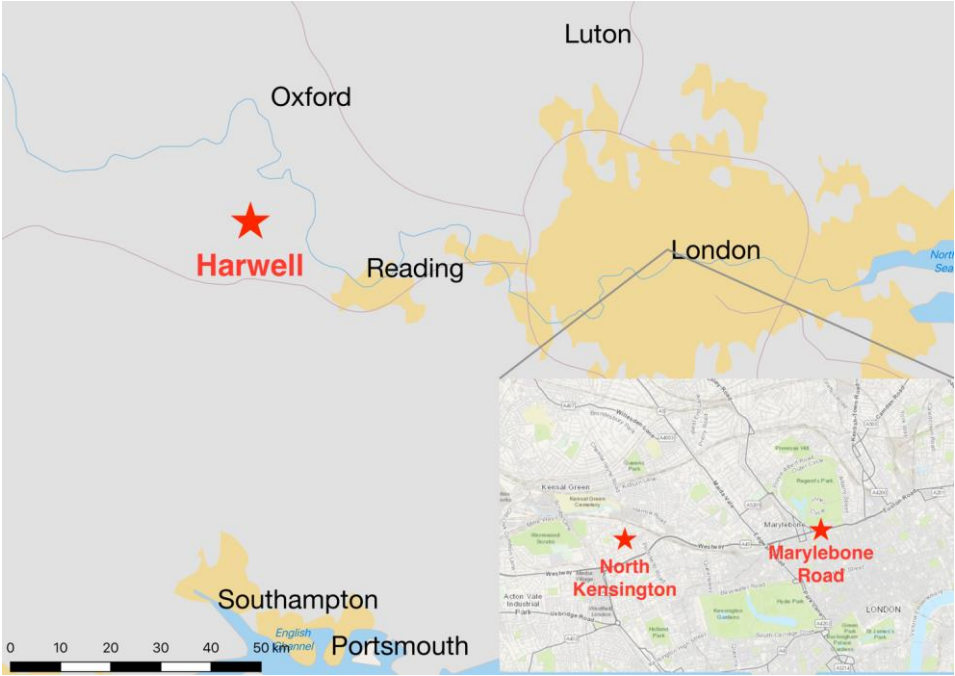
1092 \*\* Refers to events which occur simultaneously at all three sites

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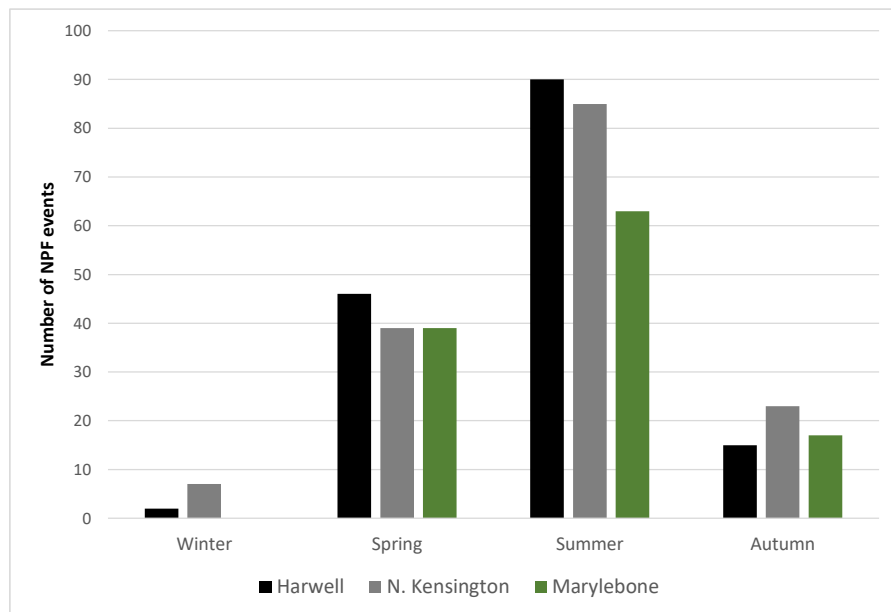
1096 **Table 2:** Annual and seasonal NSF for all areas of study.  
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	Harwell	N. Kensington	Marylebone Road
NSF <sub>NUC</sub> (Spring)	2.04±0.61	2.03±0.48	1.20±0.26
NSF <sub>NUC</sub> (Summer)	2.01±0.94	1.72±0.57	1.26±0.34
NSF <sub>NUC</sub> (Year)	2.25±0.85	1.86±0.56	1.26±0.31
NSF <sub>GEN</sub> (Spring)	1.10±0.57	1.07±0.54	1.02±0.30
NSF <sub>GEN</sub> (Summer)	1.18±0.79	1.11±0.61	1.01±0.23
NSF <sub>GEN</sub> (Year)	1.10±0.61	1.06±0.54	1.02±0.27

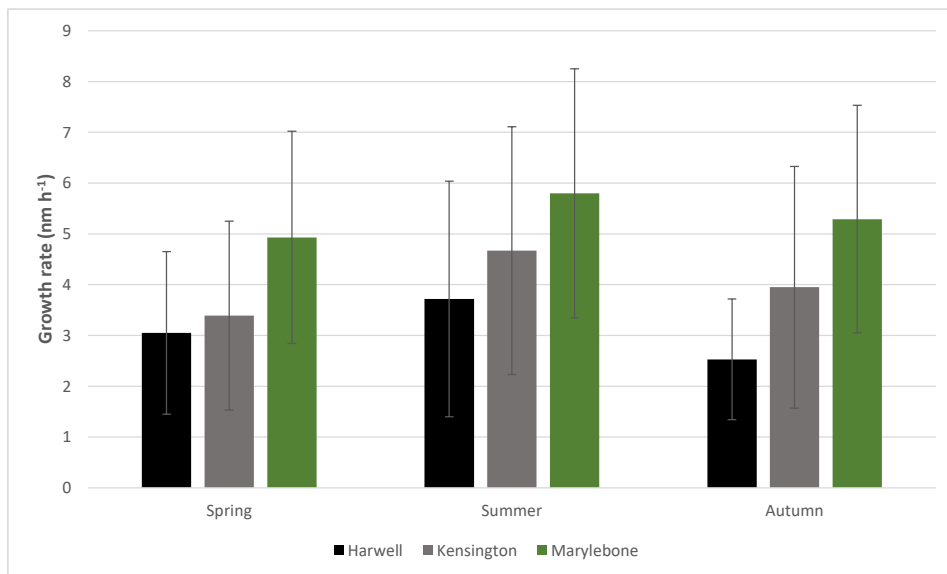
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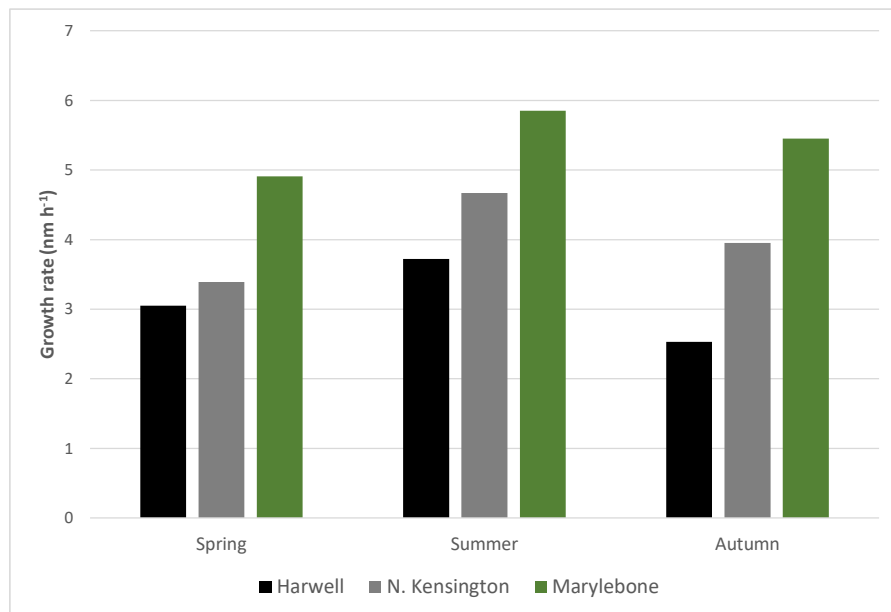


**Figure 1:** Map of the measuring stations.

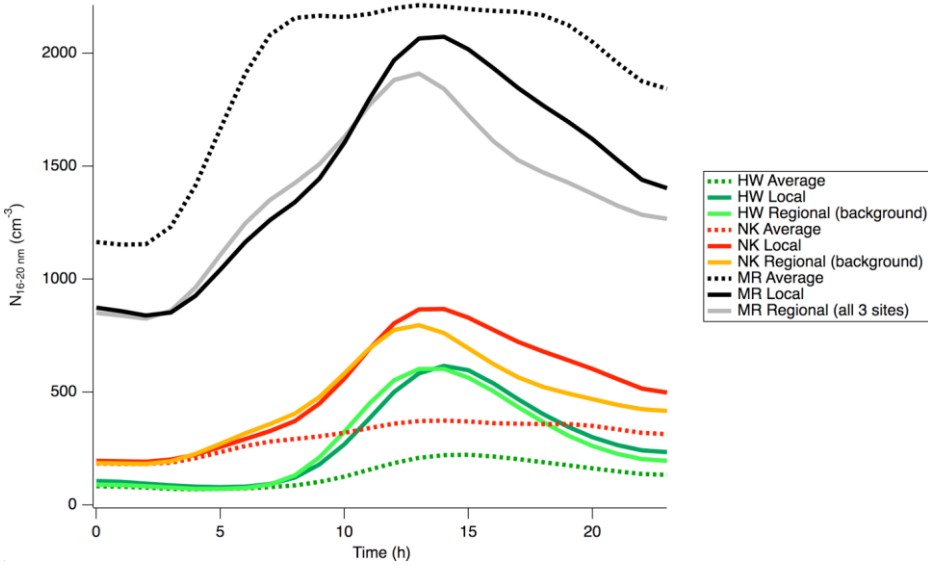


**Figure 2:** Number of NPF events per season for all seven years of the present study (Winter – DJF; Spring – MAM; Summer – JJA; Autumn – SON) at Harwell (rural), N.Kensington (urban background) and Marylebone Road (urban roadside).

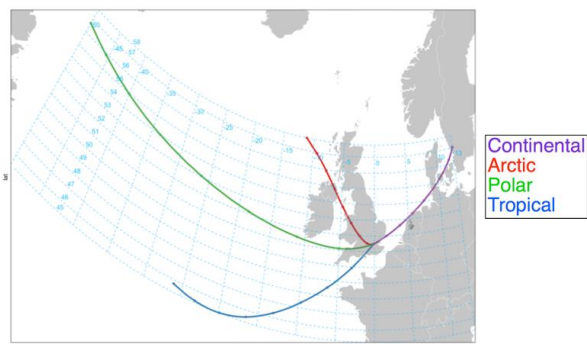




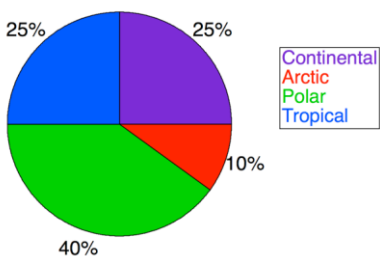
**Figure 3:** Growth rate per season at the three sites.



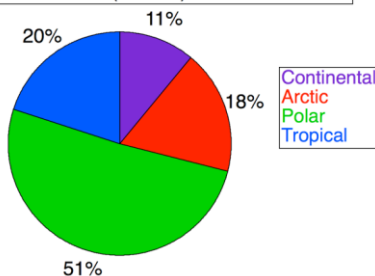
**Figure 4:** Diurnal variation of  $N_{16-20\text{mm}}$  at each site: annual [average-mean](#) and NPF event days.



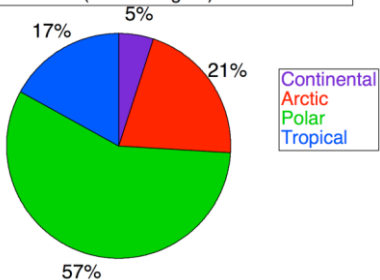
Frequency per air mass trajectory



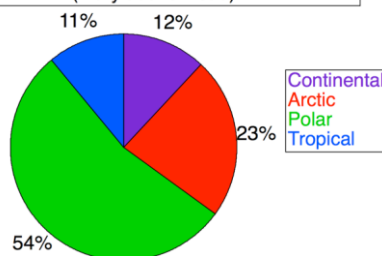
Frequency of event days per air mass trajectory (Harwell)



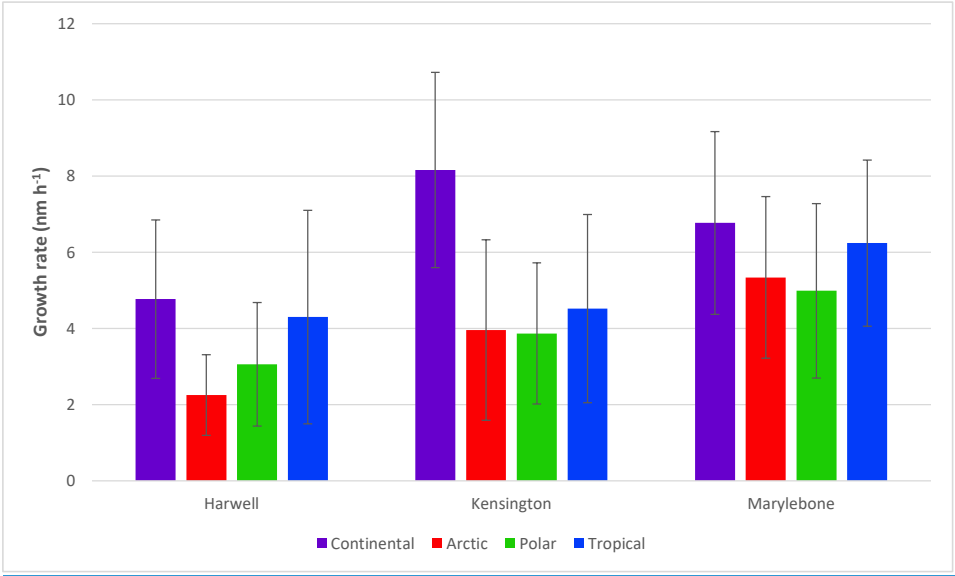
Frequency of event days per air mass trajectory (N. Kensington)

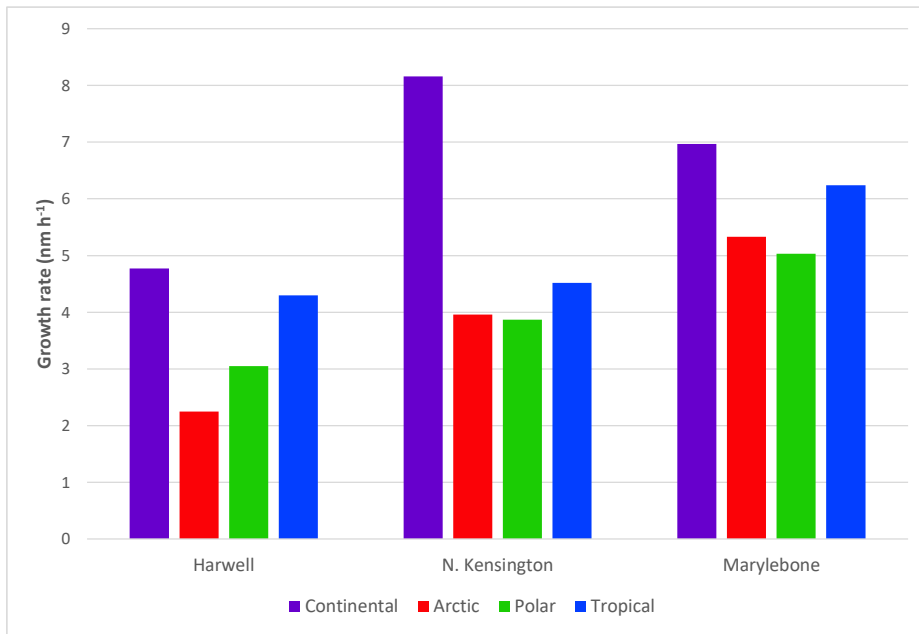


Frequency of event days per air mass trajectory (Marylebone Road)



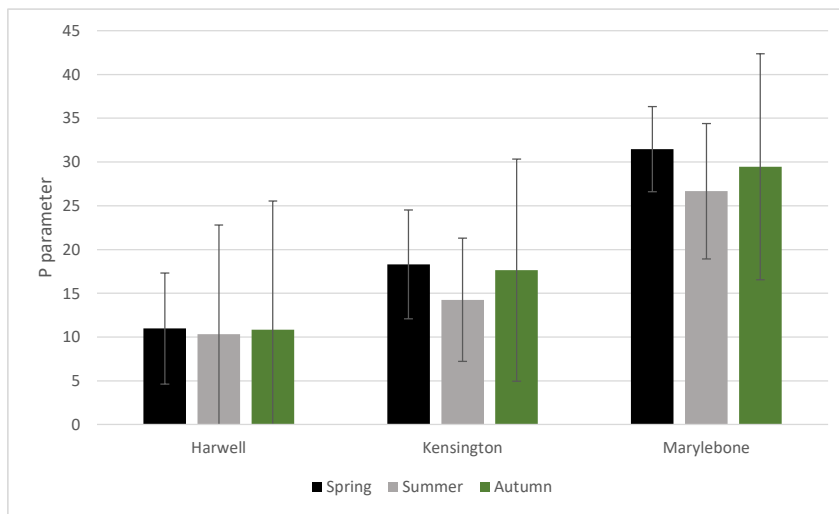
**Figure 5:** Map and frequency of incoming air mass origin – average and for NPF events per site.



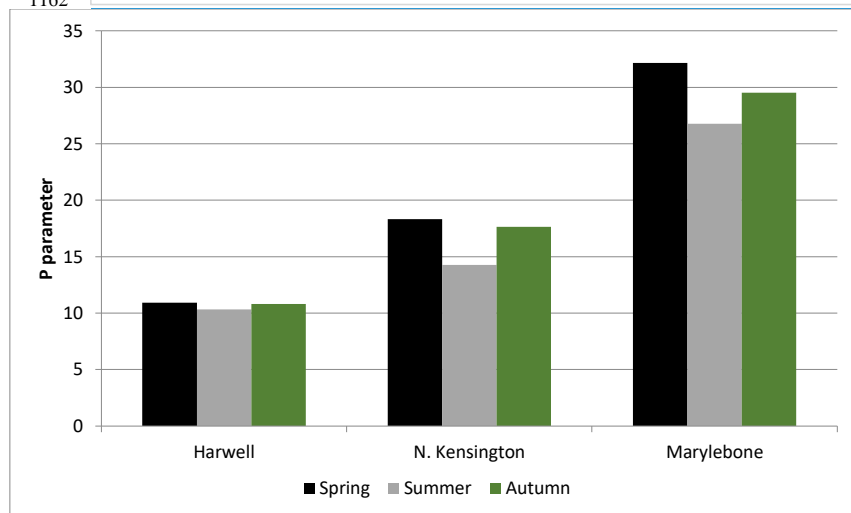


**Figure 6:** Growth rate per incoming air mass origin at each of the sites.

(a)

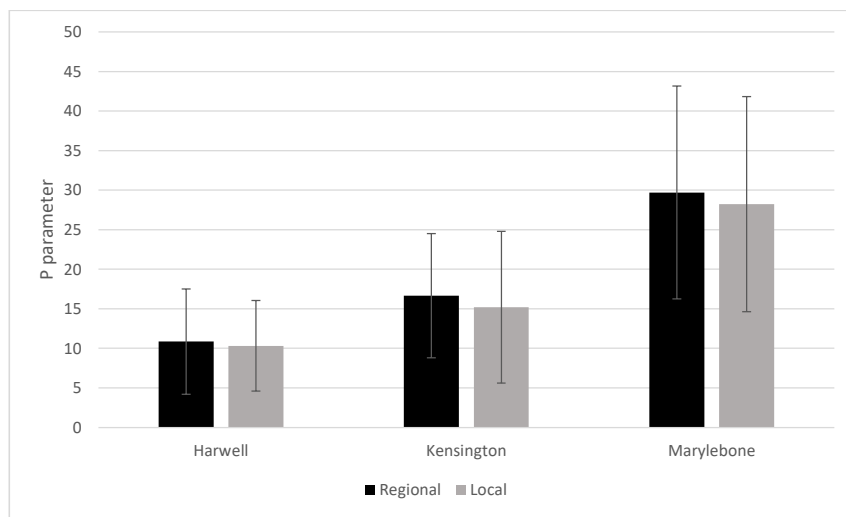


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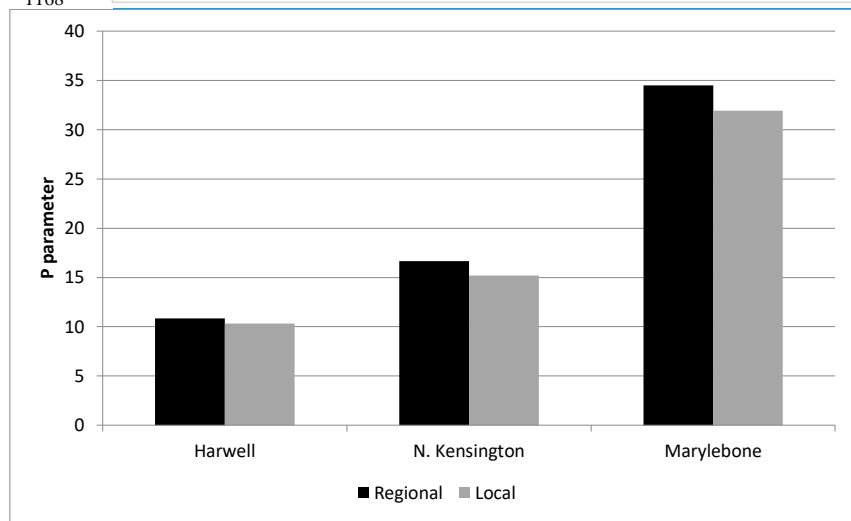


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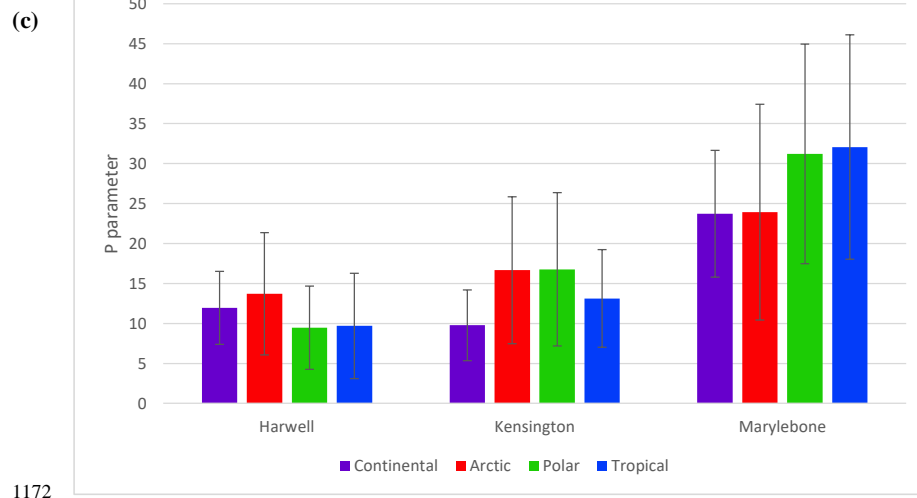
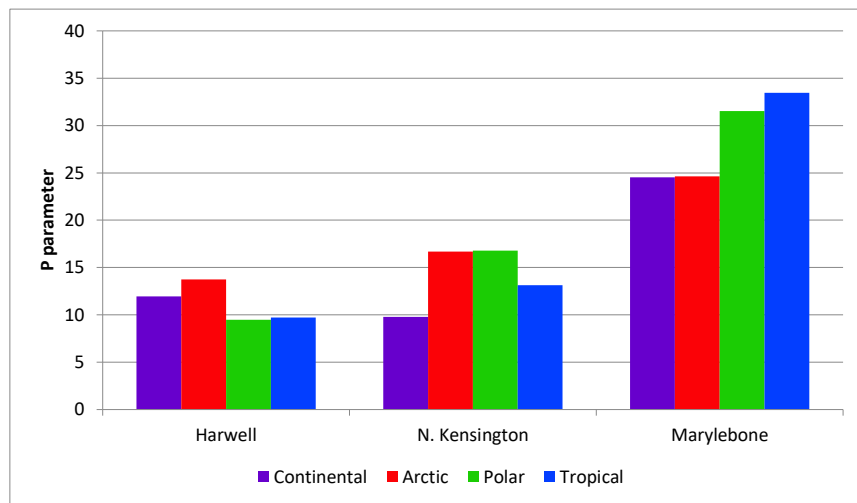
(b)



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**Figure 7:** Survival parameter P (a) per season, (b) for regional and local events (for Marylebone Road regional is for all 3 sites) and (c) by incoming air mass origin.

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