Journal: ACP - MS No.: acp-2018-1057

Title: Analysis of New Particle Formation (NPF) Events at Nearby Rural, Urban Background and Urban

Roadside Sites

Author(s): Dimitrios Bousiotis et al.

RESPONSE TO REVIEWERS

We thank the reviewers for their valuable comments, and respond point by point below.

ANONYMOUS REFEREE #2

Manuscript entitled 'Analysis of New Particle Formation (NPF) Events at Nearby Rural, Urban Background and Urban Roadside Sites' by Bousiotis et al. reports the occurrence of new particle formation events at three sites of different environments in the United Kingdom: Rural, urban background and near road sites. The authors study parameters of new particle formation such as frequency, growth rates, number concentration of sizes 16-20 nm, condensation sink, urban increment, nucleation strength factor and survival probability. The authors also report trajectory cluster analysis as well as the connection of NPF between the three different sites. In general, the manuscript, contains valuable data (three sites of different environments) and treasured statistics (7 years of data). In addition, the manuscript is well written and literature from around the world is acknowledged. However, the authors make big assumptions and conclusions without enough supporting data. The major concerns listed below need to be addressed before the manuscript is considered for publication in ACP.

Major Comments:

1. The authors report the observation of new particle formation events at three sites in the UK based on visual inspection of CPC (> 7 nm) and SMPS (> 16.6 nm). The general character of NPF events is missing. The lowest limit of the instrument is an issue and no big conclusions can be made before ensuring that the observed plume of particles is related to a new particle formation event. Authors should report how these events look like and whether they have a growing mode shape. Also, more characteristics of the growth should be reported such as possible shrinkage (see e.g. Salma et al. (2016)) and the size these particles reach. An example surface plot from each site should be added to the manuscript.

Let's take for example a regional event surface plot from Kerminen et al. (2018): figure 1, if we cannot observe the information below 16 nm, how can the authors prove that the increase in particle concentration is related to NPF, figure 2.

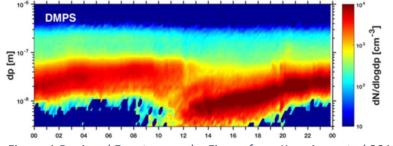


Figure 1 Regional Event example. Figure from Kerminen et al 2018.

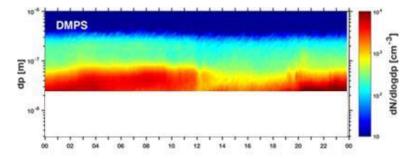


Figure 2 Modified figure 1.

The manuscript refers to many NPF studies from around the world, many of which report NPF starting from 6 nm (Salma et al., 2017), 3 nm (Dal Maso et al., 2005), 1.7 nm (Kirkby et al., 2016) while their measurement starts from 16.6 nm. The authors should present evidence that these observed particles are related to new particle formation events, and not for example a traffic growing mode (Brines et al., 2015).

RESPONSE: The dataset available, as mentioned in the text ranges from 16.6 to 604 nm. To overcome this limitation additional data was used to ensure the correct identification of the NPF events.

To achieve this:

- CPC data was used to provide insight into whether there was an increase on the number of
 particles of smaller size. An increased number of particles in the size range 7 16 nm
 (provided by the CPC data) right before or at the same time when observed in the SMPS data
 was a necessary criterion for the occurrence of an event.
- High resolution pollution data was used alongside particle number concentration data in a side by side comparison. A sharp increase in the particle number concentration which was accompanied by a similar increase in the concentrations of pollutants was an indication that these particles were probably associated with pollutant emissions. This was mainly an issue in the roadside (MR) and to a smaller extent with the background sites. Increased particle number concentrations observed at times matching the morning or evening traffic rush hours were also ignored at MR as they always coincided with increased concentrations of pollutants.
- Meteorological data was used. This mainly applies to the urban background site (NK), being in close proximity to London city centre. The possibility of a plume of pollution originating from the London city centre was considered when the site was downwind of it. A power plant to the northeast of the rural background site (HW) was also considered as a possible source of particles, though the distance is larger. Finally, as mentioned in the text, Heathrow airport and its influence were also considered.

In addition to this, the criteria set by Dal Maso et. al. (2005) were fully considered and unless there was a clear new mode of particles at the lower size range of the nucleation mode with a clear growth for at least 3 hours, an NPF event was not assigned. An example of the appearance of the events for each site has been added in the manuscript. Additionally, a discussion of particle shrinkage at later stages, which was observed at MR, is also added to the text.

Due to the limitations of the dataset, events in which the newly formed particles failed to grow to greater than 16 nm could not be seen except in the CPC data. These were rare and due to lack of additional information about their development were ignored. This clarification has been added in the text.

2. Section 2.1: Which years are studied?

RESPONSE: The years studied are 2009 – 2015. This information has been added in the text.

3. Section 2.1: Distance between the three sites should be mentioned.

RESPONSE: The distance between MR and NK is 4.5 km. The distance between HW and London city centre is about 80 km. This information has been added to the text.

- 4. Section 2.2.1: Authors report a visual inspection of CPC and SMPS data.
- How was this exactly done? Please elaborate.
- Was there any kind of counter-calibration done between these instruments?

RESPONSE: The method used was visual inspection of SMPS data supplemented by the use of CPC data to confirm the increase of the particle number concentration in the smaller size range (7 – 16 nm), as mentioned in (1). The text has been updated to clarify the method used. Both instruments are calibrated by the National Physical Laboratory according to the latest internationally recommended protocols.

- 5. Section 2.2.2: Calculation of the growth rates:
- Size of growth rates should be mentioned. E.g. growth from 7 to 20 nm? To 50 nm?

- How many points were taken in calculating the GR?
- Line 290: Authors claim that GR in NK (4.4 nm/h) are higher than the regional events GR (3.9 nm/h), what is the error bar on these calculations? Accordingly, these growth rates might be similar.

RESPONSE: As the lower size available was 16 nm, a calculation of the growth rate up to 50 nm was chosen (rather than up to 30 nm, which provided poor results in many cases due to the small range). The number of points taken depended on the development of the event and were considered from the start of the event until a) growth stopped, b) GMD reached 50 nm or c) the day ended. These points were added to the manuscript to clarify the method used.

On the third point made, due to the large variation of the growth rates of the events, the error bars are overlapped for the two groups of events. This has been included as a note in the text.

6. On line 178: the author mention nucleation mode, which is by definition number of particles between 3 and 25 nm, while the authors conduct a large study on a small fraction of this nucleation mode (16-25 nm).

RESPONSE: We are not aware of a widely recognised definition of the nucleation mode, with the term taking in different size ranges in the literature. Regardless of that, in the text it is mentioned that "NPF events are considered when a distinctly new mode of particles which appears in the size distribution at nucleation mode size, prevails for some hours and shows signs of growth", which is accurate in relation to the criteria set for NPF event selection in this study.

7. Section 2.2.4: Reference to Kulmala et al. 2017, calculating P = CS'/GR. What GR was used here? See point 4.

RESPONSE: The growth rate and condensation sink used are the ones calculated by the methods mentioned in the text. A clarification of this has been added in the text.

8. Section 3.1.1: Reference to Figure S1: cloudiness, and RH.... Is missing. Was cloudiness measured or calculated?

RESPONSE: Cloud amount data, as for all other meteorological data were measurements provided by the Met Office, as mentioned in the text. A plot with average cloud amount for each site has also been added in the supplementary.

9. Section 3.1.2: How can the authors prove that NPF events are happening at the near road site and not transported to the location?

RESPONSE: It cannot be stated with certainty whether the NPF took place at the site or particles were advected. What can be said though with confidence is that regardless of where the particle formation took place (either on the spot or in the close vicinity, as particles of that size range cannot travel to distances greater than some kilometers before either reaching detectable sizes or being diluted – especially in a polluted environment such as the London city centre), the new mode not only persists but it also grows for at least 3 hours. A clarification of this has been added in the text. If the events at the roadside site were due to advection, or a purely regional phenomenon, a much closer correlation of event days and growth rates between MR and NK would be expected than was observed.

10. The authors make big conclusions regarding the SO₂ driving mechanism of NPF which cannot be proved without adequate chemical speciation of the particles formed. These conclusions shall be minimized throughout the manuscript. Authors could try calculating sulfuric acid proxy from SO₂ and CS (Petäjä et al., 2009).

RESPONSE: In the text is stated that SO_2 was found to be lower on event days compared to the average, which logically leads to the conclusion that either the greater concentrations of SO_2 are associated with a more polluted environment with an increased condensation sink (which consequently has a negative effect in the occurrence of an event), or its concentration is adequate and it is not a factor affecting the occurrence of an event (positively or negatively). The calculation of the H_2SO_4 proxy was carried out and provided information that did not help in clarifying this point. It was found that the proxy was higher on event days at the background sites and gave an unclear result for the roadside. This result though provides no additional information as the increased values

of the proxy are the result of the higher solar radiation and the lower condensation sink found during events. Changes were made in the text to "soften" these conclusions.

ANONYMOUS REFEREE #1

The MS mainly deals with the occurrence frequency, particle growth rate, condensation sink, nucleation strength factor, survival parameter and relationships among them at 3 different locations (rural, urban background and urban roadside sites) in the UK over several years. It contains valuable results and conclusions. Some parts of the MS should be elaborated better (some items are given below as examples), and they can definitely be handled and improved. There is, however, a conceptual weakness of the study related to the lower diameter limit of the SMPS system (of 16.6 nm) which can represent the largest source of inconclusive or ambiguous interpretations for the urban sites.

Major comment:

1. New particle formation and growth events are mainly identified, separated from emission sources and classified on the basis of particle number size distributions in the particle diameter range <20 nm (e.g. Kulmala et al., Nat. Protoc., 7, 1651–1667, 2012). The diameter interval available for this in the evaluated work, namely 16.6–20 nm is quite narrow in particular, when you consider the logarithmic scale of the abscissa of size distributions. More importantly, the lower limit is requested to be even smaller (preferably below 10 nm or at 3 nm) for studies in urban atmospheric environments, where huge emission peaks can temporary dominate the smallest size ranges as well (Nieminen et al., Atmos. Chem. Phys., 18, 14737–14756, 2018). This property (16.6 nm lower limit) of the measuring system and its consequences for the data treatment, results and conclusions at the urban sites should definitely be discussed in detail, explained and resolved before any further opinion could be formed or decision can be made.

RESPONSE: The limitations and consequences due to the available dataset, as well as the additions in the method to ensure the correct selection of NPF events are explained at length in the response to Referee #2 (earlier in this document). As a result of this, clarification of the method and the additional data used have been added to the text.

Some minor comments:

1. Lines 21, 69, etc.: it is advised not to start a sentence with abbreviation.

RESPONSE: Text updated to address the comment.

2. Line 61: consider writing primary particles or emission sources instead of primary emissions.

RESPONSE: Text updated to address the comment.

3. Lines 106–109: it is unusual to attribute particles with a diameter between 1.3 and 3 nm to road traffic emissions, and, therefore, this should be discussed and explained in more detail.

RESPONSE: Text updated to accurately reflect the conclusions of the study mentioned.

4. Lines 149–151 or Table 1: supply more detailed data coverage, e.g. for each year or season of years.

RESPONSE: Table has been added in the S.I. for detailed seasonal data coverage.

5. Lines 198–203: it is requested that the diameter of particles under consideration is specified as the growth rate changes with diameter.

RESPONSE: Text updated to include the size range of the particles considered in the calculation of the growth rate.

6. Lines 262, 263, Table 2: revisit your rounding off strategy.

RESPONSE: Text and tables updated to follow a uniform rounding scheme.

7. Lines 462–463: remove; it is a repetition from lines 235–239. **RESPONSE:** Text updated to remove repeated information.

8. Fig. 2: it is unclear from the figure or related text which time interval was considered here. A number of NPF events of 90 at Harwell in summer (JJA, 92 days) should be clarified to avoid any misunderstanding. **RESPONSE:** The figure's description has updated to clarify the period plotted.

1 2 Analysis of New Particle Formation (NPF) Events at 3 Nearby Rural, Urban Background and **Urban Roadside Sites** Dimitrios Bousiotis¹, Manuel Dall'Osto², David C.S. Beddows¹, Francis D. Pope¹ and Rov M. Harrison^{1a*} 9 10 ¹ School of Geography, Earth & Environmental Sciences and 11 **National Centre for Atmospheric Science** 12 University of Birmingham, Edgbaston, Birmingham 13 **B15 2TT, United Kingdom** 14 15 ² Institute of Marine Sciences, CSIC 16 Passeig Marítim de la Barceloneta, 37-49. E-08003 17 Barcelona, Spain 18 19 20 21 22 ^aAlso at: Department of Environmental Sciences / Center of Excellence in Environmental Studies, King Abdulaziz 23 University, PO Box 80203, Jeddah, 21589, Saudi Arabia 24 25 * To whom correspondence should be addressed. 26 Tele: +44 121 414 3494; Fax: +44 121 414 3709; Email: r.m.harrison@bham.ac.uk 27

ABSTRACT

New Particle Formation (NPF) events have different patterns of development depending on the conditions of the area in which they occur. In this study, NPF events occurring at three sites of 31 differing characteristics (rural Harwell (HAR), urban background North Kensington (NK), urban roadside Marylebone Road (MR), London, UK) were studied (seven years of data). The different atmospheric conditions in each study area not only have an effect on the frequency of the events, but also affect their development. The frequency of NPF events is similar at the rural and urban background locations (about 7% of days), with a high proportion of events occurring at both sites 36 37 on the same day (45%). The frequency of NPF events at the urban roadside site is slightly less (6% of days), and higher particle growth rates (average 5.5 nm h⁻¹ at MR compared to 3.4 nm h⁻¹ and 4.2 nm h-1 at HAR and NK respectively) must result from rapid gas to particle conversion of trafficgenerated pollutants. A general pattern is found in which the condensation sink increases with the degree of pollution of the site, but this is counteracted by increased particle growth rates at the more polluted location. A key finding of this study is that the role of the urban environment leads to an increment of 20% in N_{16-20nm} in the urban background compared to that of the rural area in NPF events occurring at both sites. The relationship of the origin of incoming air masses is also considered and an association of regional events with cleaner air masses is found. Due to lower availability of condensable species, NPF events that are associated with cleaner atmospheric conditions have lower growth rates of the newly formed particles. The decisive effect of the condensation sink in the development of NPF events and the survivability of the newly formed Commented [DB(wISAS+1]: Addressing comment 1 (1st)

- 49 particles is underlined, and influences the overall contribution of NPF events to the number of
- 50 ultrafine particles in an area. The other key factor identified by this study is the important role that
- 51 urban pollution plays in new particle formation events.

52

54 1. INTRODUCTION

Ultrafine particles (particles with diameter smaller than 100 nm) typically make the greatest contribution in the total particle count, especially in urban environments (Németh et al., 2018), but a very small contribution to total volume and mass (Harrison et al., 2000). Research studies have indicated that ultrafine particles can cause pulmonary inflammation and may contribute to cardiovascular disease (Oberdörster, 2000) and have increased possibility to penetrate the brain and central nervous system (Politis, Pilinis and Lekkas et al., 2008) compared to fine and coarser particles. Since some studies report that toxicity per unit mass increases as particle size decreases (Penttinen et al., 2001; MacNee et al., 2003; Davidson et al., 2005); it is considered possible that particle number concentrations may be a better predictor of health effects than mass concentrations (Harrison et al., 2000; Atkinson et al., 2010; Kelly et al., 2012; Samoli et al., 2016). Additionally, NPF events have an impact on climate (Makkonen et al., 2012) either by increasing the number of cloud condensation nuclei (Spracklen et al., 2008; Merikanto et al., 2009; Dameto de España et al., 2017; Kalkavouras et al., 2017), or directly affecting the optical properties of the atmosphere (Seinfeld and Pandis, 2012).

69

70 The sources of ultrafine particles in urban areas can either be primary particles or emission sourcess

71 from traffic (Shi et al., 1999; Harrison et al., 2000), airports (Masiol et al., 2017) and other

e combustion related processes (Keuken et al., 2015; Kecorius et al., 2016), or by new particle

73 formation (NPF) from gaseous precursors. NPF-New particle formation as described by Kulmala et

Commented [DB(wISAS+2]: Addressing comment 2 (1st)

Commented [DB(wISAS+3]: Addressing comment 1 (1st)

al. (2014), is the process of production of low-volatility vapours, clustering of these vapours, nucleation, activation of the clusters with a second group of vapours and condensational growth to larger sizes. This process can occur both locally or on a larger scale; in the latter case the events are characterized as regional. Regional events have been found to take place in a scale of hundreds of kilometres (Németh and Salma, 2014; Shen et al., 2018), without being affected by air mass advection (Salma et al., 2016). NPF is one of the main contributors of particles in the atmosphere (Spracklen et al., 2010; Kulmala et al., 2016; Rahman et al., 2017) and this relative contribution increases moving from a kerbside to a rural area (Ma and Birmili, 2015). While NPF events in rural 81 and remote areas have been widely studied for many years (O'Dowd et al., 2002; Dal Maso et al., 82 2005; Ehn et al., 2010; Dall'Osto et al., 2017; Kalkavouras et al., 2017), in urban areas intensive studies have started mainly in recent years (Jeong et al., 2010; Minguillón et al., 2015; Peng et al., 84 2017; Németh et al., 2018). Early studies in Birmingham, UK highlighted the connection of NPF 86 events with solar radiation (Shi et al., 2001) and a low condensation sink (Alam et al., 2003), a measure of pre-existing aerosol loading (Dal Maso et al., 2002). The importance of a low 87 condensation sink was further underlined by later studies, as being one of the most influential 88 variables in the occurrence of NPF in all types of environment (Wehner et al., 2007; Park, Yum and 89 Kim, 2015; Pikridas et al., 2015). An important contributor to many NPF pathways is SO₂ (Woo et 90 al., 2001; Berndt et al., 2006; Laaksonen et al., 2008), which in the presence of solar radiation forms H₂SO₄, often the main component of the initial clusters (Kuang et al., 2008; Kulmala et al., 2013; Bianchi et al., 2016; Kirkby et al., 2016). Dall'Osto et al. (2013) pointed out that the role of SO₂ is

less significant in urban areas compared to rural and background areas. SO₂ concentration variability in urban areas was found to have a small impact on the frequency of NPF events (Alam-Shi and Harrison et al., 2003; Jeong et al., 2010), though it can have an effect on the number of particles formed (Charron, Birmili and Harrison et al., 2007). Furthermore, Dall'Osto et al. (2018) in their research at 24 sites in Europe, pointed out the different role SO₂ seems to play depending on 98 its concentration, and that of other species. Jayaratne et al. (2017) however found that in the heavily polluted environment of Beijing, China, NPF events were more probable in sulphur rich conditions rather than sulphur poor. Apart from its role in the initial formation of the clusters, H₂SO₄ seems to 101 participate in the early stages of growth of the newly formed clusters (Kulmala et al., 2005; Iida et 102 al., 2008; Xiao et al., 2015). In later stages of growth, low or extremely low volatility organic 103 compounds (O'Dowd et al., 2002; Laaksonen et al., 2008; Metzger et al., 2010; Kulmala et al., 104 2013: Tröstl et al., 2016; Dall'Osto et al., 2018) appear to be more important, while the role of 105 106 ammonium nitrate in particle growth is also considered (Zhang et al., 2017). While in rural areas the organic compounds are mainly of biogenic origin (Riccobono et al., 2014; Kirkby et al., 2016), in 107 urban areas they mainly originate from combustion procedures processes (Robinson et al., 2007; Gentner et al., 2012). Many comparative studies have reported higher growth rates in urban areas 109 compared to background sites (Wehner et al., 2007; Jeong et al., 2010; Salma, et al., 2016; Wang et 110 al., 2017), as well as greater particle formation rates (Salma, et al., 2016; Nieminen et al., 2018) and 111 a higher frequency of NPF events (Peng et al., 2017), which was attributed to the higher 112 concentration of condensable species. Salma et al. (2014) however reported fewer NPF events in the 113

city centre of Budapest compared to the urban background, due to the higher condensation sink, Due to the complexity of the conditions and mechanisms within an urban area (Harrison, 2017), NPF events are harder to study and factors to be attributed. A large number of particles of size 1.3 116 117 $\frac{3}{1}$ nm has been attributed Increased concentrations of particles in the size range 1.3-3 nm were measured at a kerbside site when being downwind from the road, following the trends of traffic-118 related nucleation mode particles, associating them to with traffic emissions at a kerbside site and 119 thus not related to resulting from homogeneous nucleation mechanisms (Rönkkö et al., 2017; 120 Hietikko et al., 2018), and studies in Barcelona, Spain (Dall'Osto et al., 2012; Brines et al., 2014) 121 122 and Leicester, U.K. (Hama et al., 2017), attributed a larger portion of nucleation mode particles to vehicular emissions compared to photochemically induced nucleation. As the condensation sink is 123 higher within an urban environment, NPF events are less favoured. Their occurrence is attributed to 124 either ineffective scavenging or the higher growth rate of the newly formed particles (Kulmala et 125 126 al., 2017), when sufficient concentrations of precursors are present in the atmosphere (Fiedler et al., 2005), as particle formation was found to take place on both event and non-event days (Riipinen et 127 al., 2007). 128 129

Commented [DB(wISAS+4]: Addressing comment 3 (1st)

7

In this study, NPF events in three areas of different land use in the southern U.K. are analyzed.

Studies for NPF events have been conducted in the past for Harwell, Oxfordshire (Charron et al.,

2007; 2008) and the effect of NPF upon particle size distributions was also considered for N.

Kensington, London (Beddows et al., 2015). A combined study including all three sites has also

130

131

132

been conducted, but in the aspect of ultrafine particle variation (Von Bismarck-Osten et al., 2013). 134 135 The present study is the first to use a combined long term database for all three sites, focusing on the trends and conditions of NPF events at these sites, as well as the first which identifies NPF 136 137 events at the highly trafficked Marylebone Road site, as up to this point ultrafine particles were attributed only to traffic (Charron and Harrison, 2003; Dall'Osto et al., 2011). As in this study a 138 rural and an urban background area are studied alongside a kerbside site in the city of London in 139 close proximity, the conditions and development of NPF events in a mid-latitude European region 140 are discussed in relation to the influence of different local environments. 141

142

143

144

2. DATA AND METHODS

2.1 Site Description and Data Availability

This study analyseds NPF events in three areas in the southern United Kingdom (Fig. 1). Harwell in 145 146 Oxfordshire, is located about 80 km west of the greater London area. The site is in the grounds of the Harwell Science Centre in Oxfordshire (51° 34' 15" N, 1° 19' 31" W) and is representative of a 147 rural background area; a detailed description of the site was given by Charron et al. (2013). North 148 Kensington is a suburban area in the western side of London, U.K, 4.5 km west of Marylebone 149 Road. The site is located in the grounds of Sion Manning School (51° 31' 15" N, 0° 12' 48" W) and 150 is representative of the urban background of London. A detailed description of the site was given by 151 Bigi and Harrison (2010). Marylebone Road is located in the centre of London, U.K. The site is 152 located on the kerbside of Marylebone road (51° 31' 21" N; 0° 9' 16" W), a very busy arterial route 153

Commented [DB(wISAS+5]: Addressing comment 3 (2nd)

within a street canyon. A more detailed description of the area can be found in Charron and Harrison (2003).

156

At all three sites, seven years (2009 - 2015) of particle number size distributions in the range of 157 16.6 – 604 nm have been measured and recorded as 15-minute averages, using a Scanning Mobility 158 Particle Sizer (SMPS), comprised by an Electrostatic Classifier (EC, TSI model 3080) and a 159 condensation Particle Counter (CPC, TSI Model 3775), operated on behalf of the Department for 160 Environment, Food and Rural Affairs (DEFRA) in the U.K. At all sites the inlet air is dried, and 161 operation is in accord with the EUSAAR/ACTRIS protocol (Wiedensohler et al., 2012). These 15-162 minute measurements were averaged to an hourly resolution. In Harwell there were 46930 hours of 163 164 available SMPS data (76.5% coverage), in N. Kensington 51059 (83.3% coverage) and in at Marylebone Road 45562 (74.3% coverage). Detailed data availability is found in *Table S1. - A free-165 166 standing CPC (TSI model 3022A) also operated alongside for most of the years of the survey and was used to give an estimate of particles in the 7-16.6 nm range by difference from the SMPS. 167

168

Additionally, air pollutants and other aerosol chemical composition data were extracted from the
DEFRA website (https://uk-air.defra.gov.uk/). Meteorological data for Harwell and Heathrow
airport (used for N. Kensington and Marylebone road) were available from the Met Office, while
solar radiation data from Benson station (for Harwell) and Heathrow airport (for N. Kensington and
Marylebone Road), were extracted from the Centre for Environmental Data Analysis (CEDA) site

Commented [DB(wISAS+6]: Addressing comment 2 (2nd)

Commented [DB(wISAS+7]: Addressing comment 4 (1st)

(http://www.ceda.ac.uk). Back trajectory data calculated using the HYSPLIT model (Draxler and Hess, 1998), were extracted by the NOAA Air Resources Laboratory (https://ready.arl.noaa.gov/READYtransp.php) and were processed using the Openair package for R (Carslaw and Ropkins, 2012).

178

179

2.2 Methods

180 2.2.1 NPF events selection

The identification of the NPF event days was made by visual inspection of SMPS data. 181 182 supplemented with the use of and CPC data to confirm the formation of a new mode of particles, using the criteria set by Dal Maso et al. (2005). NPF events are considered when a distinctly new 183 184 mode of particles which appears in the size distribution at nucleation mode size, prevails for some hours and shows signs of growth. Using these criteria, NPF events are classified into two classes, I 185 186 and II depending on the confidence level. Class I events are further classified to Ia and Ib, with class Ia containing very clear and strong particle formation events, while Ib contains less clear events. In 187 this study the events of class Ia are only are considered as being the most suitable for analysing case 188 studies of NPF events (Figure S1). At this point it should be mentioned that due to the particle size 189 190 range available, NPF events in which new formed particles failed to grow beyond 16.6 nm (if any) could not be identified. Though such rare formations occasions were identified using the CPC data, 191 bursts of new particles in the size range < 16.6 nm that did not appear on the SMPS dataset were 192 193 ignored as their development was unknown. High time resolution data for gaseous pollutants and

Commented [DB(wISAS+8]: Addressing comment 4 (2nd)

Commented [DB(wISAS+9]: Addressing major comment 1 (1st) and 1 (2nd)

aerosol constituents was used to identify pollution events affecting particle concentrations and these
were removed from the data analysis. –This analysis took account of the fact that nanoparticle
emissions from Heathrow Airport affect size distributions at London sites (Harrison et al., 2018),
and such primary emission influences were not included as NPF events.

2.2.2 Calculation of the condensation sink and growth rate

For the calculation of the condensation sink the method proposed in Kulmala et al. (2001) was used

201 in which the condensation sink is calculated as

$$\mathbf{CS} = \mathbf{4\pi D} \sum_{\mathbf{\beta_M}} \mathbf{r} \mathbf{N}$$
203

where r is the radius of the particles and N is the number concentration of the particles. D is the diffusion coefficient calculated (for T = 293 K and P = 1013.25 mbar) according to Polling et al. (2000):

$$D_{vap} = 0.00143 \cdot T^{1}.75 \frac{\sqrt{M_{air}^{-1} + M_{vap}^{-1}}}{P\left(D_{x,air}^{\frac{1}{3}} + D_{x,vap}^{\frac{1}{3}}\right)^{2}}$$
(2)

where P is air pressure, M is the molar mass and D_x is the diffusion volume for air and H_2SO_4 . β_M is

the Fuchs correction factor calculated as (Fuchs et al., 1971):

$$\beta_{\rm M} = \frac{1 + K_{\rm n}}{1 + \left(\frac{4}{3a} + 0.377\right) K_{\rm n} + \frac{4}{3a} K_{\rm n}^2} \tag{3}$$

where K_n is the relation of the particle diameter and the mean free path of the gas λ_m , called the

Knudsen number.

The growth rate of the particles on nucleation event days was also calculated as proposed by

Kulmala et al. 2012, using the formula

$$GR = \frac{D_{P_2} - D_{P_1}}{t_2 - t_1}$$
 (4)

for the size range 16.6 - 50 nm. The number of points taken depended on the development of the

event and were considered from the start of the event until a) growth stopped, b) GMD reached 50

nm or c) the day ended.

for the period of each event day when growth was observed.

Commented [DB(wISAS+10]: Addressing comments 5 (1st)

2.2.3 Calculation of the urban increment (U.I.)

- 231 The urban increment is defined as the ratio of the number concentration of particles below 20 nm
- 232 for event days to the average (for the period April October, when the majority of the events take
- 233 place) for North Kensington to that at Harwell. This provides with a measure of the new particles
- formed in each area in comparison to the average conditions, and is calculated by

235

$$\frac{NK_{Nuc Max} - NK_{Bg}}{HW_{Nuc Max} - HW_{Bg}}$$
236 U.I. =
$$\frac{WK_{Nuc Max} - W_{Bg}}{HW_{Nuc Max} - HW_{Bg}}$$
(5)

237

- where $NK_{Nuc\ Max}$ is the maximum concentration of particles below 20 nm found in the diurnal cycle
- 239 on event days (found at 13:00) and NK_{Bg} is the average mean concentration at the same time (same
- 240 for Harwell in the denominator).

241

242 2.2.4 Calculation of nucleation strength factor (NSF) and the P parameter

- 243 The Nucleation Strength Factor (NSF) was proposed as a measure of the effect nucleation events
- 244 have in the composition of ultrafine particles in an area. Two factors were proposed. First is the
- 245 NSF_{NUC}. This is calculated as

246

$$NSF_{NUC} = \frac{\left(\frac{N_{(smallest size available-100)}}{N_{(100-largest size available)}}\right)_{nucleation days}}{\left(\frac{N_{(smallest size available-100)}}{N_{(100-largest size available)}}\right)_{non-nucleation days}}$$
(6)

and provides of a measure of the concentration increment on nucleation days exclusively caused by
new particle formation (NPF). The second factor is NSF_{GEN} calculated as

$$NSF_{GEN} = \frac{\left(\frac{N_{smallest\ size\ available} - 100}{N_{100-largest\ size\ available}}\right)_{all\ days}}{\left(\frac{N_{smallest\ size\ available} - 100}{N_{100-largest\ size\ available}}\right)_{non-nucleation\ days}}$$
(7)

and gives a measure of the overall contribution of NPF on a longer span (Salma et al. 2017).

254 The dimensionless survival parameter P, as proposed in Kulmala et al. (2017), was calculated as

 $\mathbf{P} = \frac{\mathbf{CS'}}{\mathbf{GR'}}$

where $CS' = CS/(10^{-4} \text{ s}^{-1})$ and $GR' = GR/(1 \text{ nm hour}^{-1})$. CS and GR values used as were calculated with the methods mentioned at 2.2.2. An increased P parameter is an indication that a smaller percentage of newly formed particles will survive to greater sizes. Hence this is the inverse of particle survivability, and values of P<50 are typically required for NPF in clean or moderately polluted environments, although higher values of P are observed in highly polluted atmospheres (Kulmala et al, 2017).

3. RESULTS AND DISCUSSION

266

267 3.1 NPF Events at in the Background Areas

268 3.1.1 Conditions and trends of NPF events

269 The number of NPF event days for each site per year, those that took place simultaneously on-at both urban and rural background sites, as well as those events that took place in at all three sites 270 simultaneously appear in Table 1. Given that overall data recovery was in the range of 74-83%, 271 results from individual years are unreliable, but the seven-year runs should average out most of the effects of incomplete data recovery. The number of events is similar for Harwell and N. 273 274 Kensington, with a frequency of about 7% of all days with data. There is a clear seasonal variation favouring summer and spring (Figure 2) for both areas of the study. A similar pattern of variation 275 was found for N. Kensington by Beddows et al. (2015). In general, higher solar radiation, lower 276 relative humidity, low cloud cover and higher pressure conditions, lower concentrations of 277 278 pollutants as well as lower condensation sink are found when NPF events took place in all areas (Figure S2+), as was also reported by Charron et al. (2007) for Harwell. While SO2 is one of the 279 main factors for NPF events to occur, concentrations are lower when events take place. This is 280 indicative that SO₂ concentrations in these areas are sufficient for events to take place, and higher 281 282 concentrations are likely to be associated with higher pollution and a higher condensation sink. The proxy for [H₂SO₄] was calculated for the background sites using the method outlined in (Petäjä et 283 al., (2009) and was found to be higher on event days for both background sites (results not 284 included). This indicates the possible positive effect of increased concentrations of H₂SO₄ in the 285

Field Code Changed

287 of either the solar radiation (via the formation of OH radical) or the reduced condensation sink to its formation. This is also Similar is For the case for of gaseous ammonia (results not included) for 288 289 Harwell where data was available, as there was no distinct variation found between event and nonevent days, but as the concentration of ammonia in the U.K. is in the range of few ppb (Sutton et al., 290 1995), it is sufficient according to ternary nucleation theory (Korhonen et al., 1999) for NPF events 291 not to be limited by ammonia. The average growth rate for Harwell was found to be 3.37-4 nm h⁻¹, 292 within the range given by Charron et al. (2007) and higher at N. Kensington at 4.22 nm h⁻¹, a trend 293 294 found for all seasons (Figure 3). The increased growth rate in the urban area can be related to the greater presence of organic matter and other condensable species. In both areas NPF events had 295 higher growth rates in summer than in spring, as was also found in previous studies (Kulmala et al., 2004: Nieminen et al., 2018). This may be associated with the higher presence concentration of 297 298 organic compounds emitted by trees during summer (Riipinen et al., 2007), or faster oxidation rates due to higher concentrations of hydroxyl radical and ozone (Harrison et al., 2006). 299 300 About 45% of the events took place simultaneously in both background areas. These events are 301

occurrence of NPF events as well as, since SO₂ concentrations were found lower, the increased role

Commented [DB(wISAS+11]: Trying to address comment 10

Commented [DB(wISAS+12]: Addressing comment 6 (1st)

characterized as regional, as NPF takes took place in on a larger scale, regardless of the local

conditions on of the given area. In this case, meteorological conditions were even clearer, indicative

of the greater dependence of regional events on synoptic conditions rather than local. While most

chemical constituents were also lower in concentration during regional events, different patterns

302

303

were found for organic compounds and sulphate for each background area. In Harwell sulphate was
higher during regional events, while in N. Kensington organic compounds were higher during
regional events. This may be indicative of the variable role that specific chemical species have in
condensational nanoparticle growth (Yue et al., 2010). In all cases though, the concentrations of
these species were lower compared to the average conditions. Despite these differences, the growth
rate of particles was found to be higher for local events in N. Kensington (4.4 nm h⁻¹) compared to
regional events (3.9 nm h⁻¹), though within the margin of error uncertainty. In Harwell, no difference

Commented [DB(wISAS+13]: Addressing comment 5 (2nd

was found in the growth rate between regional and local events.

314

315

313

3.1.2 Variability of the origin of the air masses on NPF events

As both background areas sites are relatively close to each other (about 80 km) and had similar number of event days, a combined clustering of back trajectories for the event days (only) in these two areas was attempted. This would provide an insight into the origin of air masses for local and regional events, as well as the conditions for these air masses. The data for local N. Kensington events and both local and regional events in Harwell were clustered together and the results along with the characteristics of the air mass clusters are found in Figure S32.

322

Cluster C3, which is placed between C2 and C4 among those originating from the Atlantic Ocean, has the highest percentage for both area specific and regional events. Specifically, for regional events the percentage is over 35%, much higher compared to all other, showing a clear "preference"

of regional events for cleaner and faster moving air masses from mid-latitudes of the Atlantic 326 327 Ocean. This "preference" explains the lower production and growth rate of the new particles found for regional events, compared to local ones, as air masses from this area have lower organic carbon 328 329 and SO₂ concentrations. Cluster C5, originating straight from the north but representing air masses that have crossed the Irish Sea and have not extensively gone over land presents a similar case. 330 These cold and clean air masses are associated with a low growth rate and survivability of the 331 newly formed particles. Local events for both sites apart from those in Cluster C3 are highly 332 associated with Clusters C1 and C2, C1, which contains slow and polluted air masses, presents the 333 334 highest growth rate and as a result high particle survivability, as given by the P parameter (see later). On the other hand, C2 which consists of warm and moist air masses from lower latitudes is 335 the least common for regional events and presents high growth rate and survival probability of the 336 particles. Apart from the weak relation found with particulate organic carbon concentrations and 337 338 growth rate (Figure S32), there appears to be an inverse relation between the temperature and survivability of the particles. Warmer air masses seem to be related to higher particle survival 339 probability, which may be attributable to greater growth rates as temperature increases (Yli-Juuti et 340 al., 2011). 341

342

343 3.1.3 Urban increment and particle development

The urban environment, depending on the conditions, may have a positive or negative effect in the number of the particles formed and their consequent survival and growth. Both Harwell and N. Kensington are in background areas, rural and urban respectively. As a result, while the concentrations of pollutants are higher in N. Kensington than Harwell, their effect is smaller compared to that of Marylebone Road. A comparison of the particles smaller than 20 nm, gives insight into the formation and survival of the newly formed particles in the initial stages. Calculating the urban increment (equation 5) using the two background sites showed around 20% more particles of size 16 - 20 nm in N. Kensington than Harwell for event days, an increment that is even stronger when solely local events are considered (Figure 4). As the sizes of the particles in the calculation are relatively large and due to the higher condensation sink found in N. Kensington, this increment is expected to be larger for smaller size particles. A possible explanation for this result may be the greater concentration of organic compounds which is observed in N. Kensington, as discussed earlier, which leads to more rapid formation of secondary condensable species that enhances the nucleation process in the more polluted area.

Considering the local events, most of the pollutant <u>concentration</u> data available appear to be higher which is reflected in the condensation sink as well. The role of the polluted background appears to be decisive in the further growth of the newly formed particles, especially for Harwell. This, at both sites causes the number of particles of greater size to be smaller for the later hours in the days of local events (Figure S43). Another possible reason for this difference in the larger size ranges can be the higher concentration of organic content on the days of regional events at N. Kensington (as discussed earlier). On the other hand, for Harwell all hydrocarbons with available data are lower

throughout the day (apart from ethane) during regional events. Unlike N. Kensington, at Harwell particles smaller than 20 nm as well as the growth rate of the newly formed particles are almost the same for regional and local events.

The calculation of the increment in Marylebone Road provided negative results; particles smaller than 20 nm were less abundant on event days compared to the average, throughout the day. This is due to the fact that Marylebone road is heavily affected by traffic pollution and on average, conditions do not promote NPF events due to the high condensation sink, unless clear conditions prevail, which are also associated with a low particle load.

3.2 NPF Events at Marylebone Road

For many years, NPF events were thought not to take place in heavily polluted urban areas, as the effect of the increased condensation sink was considered detrimental crucial in suppressing the formation and growth of new particles. Recent long term analyses have shown this is not the case and nowadays an increasing number of studies studies confirm the occurrence of NPF events in urban areas. In this study, for the same period of seven years as for the two background areas, NPF events were found to occur for 6.1% of days at Marylebone Road, lower than in the background areas. Though, due to the particle size range available there cannot be a definitive answer to whether the formation of the particles takes place in the specific arealocality of the sampling site, due to the observed increase in particle concentrations in the range 7 – 16 nm (provided by the CPC data) and the increased growth rates found in urban areas in general, it can be assumed that the
formation takes place either in the area of the measuring site or in its close vicinity, while the
growth of the particles persists in the area for several hours, despite the high condensation sink.
Seasonal variation is similar to that at the background sites, but day of the week variation is stronger
at Marylebone Road further favouring weekends (Figure S54), as on these days traffic intensity is

391

392

lower.

Commented [DB(wISAS+14]: Addressing comment 9 (2nd)

In general, similar conditions found in the background areas to affect NPF events at the background 393 394 sites are also found at Marylebone Road, despite a much larger condensation sink. (Figure S21). As a result, less particles of size smaller than 20 nm were found on NPF event days than the average 395 396 for the site, as the sum of background particles plus those formed on these days were less than that on an average day. The growth rate of the newly formed particles (5.5 nm h⁻¹), is higher than that of 397 the background sites (5.5 nm h⁺), which is in agreement with the findings in the study of the 398 background areas on the possible role of the condensable species, the concentrations of which are 399 even greater at the urban kerbside. About 15% of NPF event days at Marylebone Road presented 400 401 particle shrinkage after the initial growth; the study of these cases though is outside of the context of the present work. At Marylebone Road, the number of NPF days which were common with the 402 background sites was fewer, as local conditions (high condensation sink) are detrimental to the 403 occurrence of NPF events and thus the days of regional events including Marylebone Road were 404 separately studied for this site. The regional event days that were common for all three sites were 37 405

Commented [DB(wISAS+15]: Addressing comment 1 (2nd)

406 (31% of events at Marylebone Road) (Table 1). As with the other two areas, the growth rate is 407 higher during local events, but the conditions are mixed, with lower concentrations of sulphate and 408 organic compounds but higher SO₂, NOx and elemental carbon. The relationship with higher wind 409 speed (mainly western) (Figure So₂), solar radiation (which results in greater H₂SO₄ formation) and 410 lower relative humidity, indicate the stronger relation of the regional events with synoptic 411 conditions than the local events in the heavily polluted environment of Marylebone Road.

412

413

414

3.3 Connection of NPF Events with Incoming Air Masses

3.3.1 Air mass back trajectory clustering and connection with NPF events

415 The origin of the air masses plays a very important role in the occurrence of NPF events, as shown in Section 3.1.2. Air masses of different origins have different characteristics. Back trajectories 416 provide excellent insight into the source of the air masses. Air mass back trajectories were 417 418 calculated both for all days and for NPF event days for each site separately, with the aim of complementing the analysis in Section 3.1.2 which addressed only the event days. The additional 419 analysis gives a view of the frequency of NPF events within different air mass types. The initial air 420 mass back trajectory clustering ended up with an optimal solution of 9 clusters of different air 421 masses. As many of these clusters had similar characteristics and origin, solutions with fewer 422 clusters were attempted. As the number of clusters was decreasing clusters became a mixture of 423 different origins, thus making the distinction of different sources harder. As a result, the method 424

- 425 chosen was to merge clusters of similar origin and characteristics, which kept the detail of the large
- 426 number of clusters and made the separation of the different origins more distinct.

427

435

439

- 428 The resulting four merged clusters (Figure 5), using the characterisation proposed by McIntosh et
- 429 al., (1969), are:
- An Arctic cluster, which originates mainly from the northerly sector. It occurs about 10% of
- 431 the time and consists of cold air masses, which either passed over northern parts of the U.K. or
- through the Irish Sea.
- 433 A Tropical cluster, which originates from the central Atlantic. It occurs 25% of the time and
- 434 contains warmer air masses. A small percentage of this cluster contains masses that have
 - passed over countries south of the U.K. Even though these days were more polluted, the
- distering method was unable to clearly distinguish these days as it does not take into account
- particle numbers or composition, even when the 9-cluster solution was applied.
- A Polar cluster, which originates from the north Atlantic. It is the most common type of air
 - mass arriving in the areas of study and occurs about 40% of the time bringing fast moving,
- "clean" air masses with increased marine components (Cl, Na, Mg) from the west. This cluster
 - also contains airmasses that have passed through Ireland, though an effect on particle size and
- chemical composition is not distinct.
- A Continental cluster, which originates from the east. It occurs about 25% of the time and
- 444 consists mainly of slow moving air masses, originating from the London area (for the

background areas) and/or continental Europe. It has higher concentrations of most pollutants as well as the highest condensation sink.

The occurrence of each air mass class for average and event days for Harwell and London (both sites) can also be found in Figure 5, while their main characteristics for each site can be found in Table S24. Though in this case the air mass grouping for each site was done in a different analysis, the resulting groups are almost identical in their characteristics and frequency, as the sites are close to each other.

The Polar cluster is the one prevailing on both average and event days. This consists of clean fast-moving air masses originating mainly from mid and high latitudes of the Atlantic, and this cluster presents favourable conditions for NPF events. The association of NPF events with air masses from the mid-Atlantic at N. Kensington was also found by Beddows et al. (2015). Cool Arctic air masses on average are not clean as they may have passed over the northern U.K. The event days associated with this air mass type have the lowest concentrations of the pollutants within available data for all areas. The increased percentage of events with this air mass at all sites indicates that lower temperatures, in a clear atmosphere with sufficient solar radiation are favourable for NPF events as found in previous studies (Napari et al., 2002; Jeong et al., 2010; Kirkby et al., 2011). A similar trend of increased probability with polar and arctic maritime air masses was also found for Hyytiälä, Finland by Nilsson et al. (2001). Tropical air masses have a lower probability for NPF events,

which is associated with the fact that a number of these days are associated with air masses which 465 have passed from continental areas south of the U.K. (France, Spain etc.). Specifically for 466 Marylebone Road the NPF probability is a lot lower (11% versus 17% for N. Kensington and 20% 467 468 for Harwell). This is due to the fact that these air masses are more related to southerly winds which, in Marylebone Road are associated with a street canyon vortex which causes higher pollutant 469 concentrations at this site. Finally, the Continental cluster presents the lowest probability for NPF 470 events. The air masses in this group originate from continental Europe and for the background areas in most cases have passed over the London region as well. This results in both a higher 472 condensation sink and concentration of pollutants, which limits the number of days with favourable 473 conditions for NPF events. Growth rate for all sites though appears to be higher for air masses 474 475 originating from more polluted areas (Figure 6), which appear to enhance the growth process due to containing a higher concentration of condensable species (after oxidation). 476

477

478 3.4 Nucleation Strength Factor (NSF)

The NSF (equations 6 and 7) is used to describe the effect nucleation events have on the number of particles at a site. The values of NSF for each site and for seasons spring and summer are shown in Table 2. The decrease of the contribution of NPF events to particle number, moving from the rural area to the kerbside was also found in previous studies (Salma et al., 2014; 2017). This is explained by the increased contribution to the particle number concentrations of other sources, mainly combustion in the urban environment, compared to rural areas. Apart from this trend, in the

background areas the increase of N_{16-100} was greater in spring than summer. This effect seems stronger in the urban background area compared to the rural, as in that area the variability of N_{16-100} is greater for event days compared to that of the rural area. On the other hand, the contribution of NPF events in the longer span, as is illustrated by the NSF_{GEN} appears to favour summer for all areas, showing the increased formation and survivability of particles in this season.

490

485

486

487

488

489

For Marylebone Road the result for the increase of the N_{16-100} is greater in summer than in spring, in 491 contrast to what was found for the background sites. This is due to the fact that in summer the 492 493 traffic intensity is decreased, giving the contribution from NPF events a stronger effect compared to the other sources. The very small increase found on NPF events in Marylebone Road, with a factor 494 of just 1.26, a lot lower than that found in the urban area of Seoul, South Korea (Park et al., 2015), 495 is indicative of the reduced effect of NPF events in an area which is heavily affected by traffic, as 496 497 also pointed out by Von Bismarck-Osten et al. (2013) in their study on particle composition in Marylebone Road. 498

499

500

501

504

3.5 The Survival Parameter P

detectable sizes. The average values of the P parameter for each of the areas of this study are 10.5 for Harwell, 15.8 for N. Kensington and 28.9 for Marylebone Road. The values found put

The survival parameter P is a measure of the probability for newly formed particles to survive to

Commented [DB(wISAS+16]:

Commented [DB(wISAS+17]: Deleted text to address comment 7 (1st)

Marylebone Road to the upper end of heavily polluted areas in Europe, North Kensington to the

same level as many other urban areas in Europe, while Harwell had somehow higher values 505 compared to other rural background areas in Europe, as calculated by Kulmala et al. (2017). The 506 seasonal, air mass origin and local versus regional variations can be found in Figure 7 (winter is 507 508 excluded due to very low number of events). While the increasing trend of the P parameter as we move from rural background to kerbside was expected, it can be seen that there is a clear seasonal 509 pattern in all three areas, with summer having the lowest P parameter (greatest survivability) 510 compared to the other two seasons. This is associated with the higher growth rate found in summer 511 for all areas of this study, as the differences in the condensation sink on event days are negligible 512 between seasons. The case is similar for regional and local events. The result per air mass origin is 513 related to the different conditions and parameters of each incoming air mass in each area. For 514 example, the higher P parameter for Tropical air masses at Marylebone Road, is associated with the 515 higher condensation sink found for this kind of air masses, due to the street canyon effect which is 516 517 specific for Marylebone Road for southerly wind directions with which these air masses are mainly related, while the higher values for the rather clean Arctic air masses for the other two areas are 518 associated with the lower growth rates found for this kind of air mass in these areas. The more 519 polluted Continental air masses seem to have a different effect for rural and urban areas. Their 520 higher condensation sinks and concentrations of pollutants have a negative effect on P-values for 521 the rural site and a positive effect at the urban sites. The exact opposite is found for the cleaner air 522 masses of the Polar cluster, which appear to result in reduced P-values of the newly formed 523

particles at the urban sites. This is related to the lower condensation sink associated with this air mass type.

526

527

543

4. CONCLUSIONS

Seven years of data from three distinct areas (regional background, urban background, kerbside) in 528 the southern U.K. were analysed and the conditions associated with NPF events were studied. NPF 529 events were found to occur on about 7% of days at background sites and less at the kerbside site. 530 The conditions on event days for all three areas were similar, with clear atmospheric conditions and 531 532 a lower condensation sink. While the condensation sink appears to be the most important factor limiting NPF events at the kerbside site, SO2 was found to have smaller concentrations on event 533 534 days for all areas, which indicates that either on average it is in sufficient concentrations for NPF events to occur, or that other variables that participate in the production mechanism of H₂SO₄ are 535 536 more important. The growth rate of the newly formed particles increases from the rural site to the kerbside and is greater in summer compared to other seasons for all three sites. Almost half of the 537 NPF events at the rural and urban background sites were found to happen simultaneously. In these 538 cases, the atmospheric conditions were cleaner, which resulted in slower growth rates. While most 539 of the chemical species available were at lower concentrations in regional events, a difference in the 540 behaviour with respect to sulphate and organic compounds was found between the two background 541 site types. 542

Commented [DB(wISAS+18]: Trying to address comment 10

The prevailing origin of air masses in the southern U.K. is from mid and high latitudes of the Atlantic Ocean. These fast-moving air masses present an increased probability for NPF to occur. The case is similar for the cooler and cleaner arctic air masses, while air masses from the tropics and continental Europe, having greater pollutant content, have decreased NPF probability, but a higher growth rate of particles when NPF events occurred. Regional events appear to be more associated with cleaner air masses, presenting a smaller growth rate and condensation sink compared to local events. The difference in growth rate is probably related to the greater content of condensable species; a positive relation of particle survival probability with temperature was also found.

Comparing the background areas in this study, particles of 16-20 nm were found to be about 20% greater in concentration (above long-term average) on NPF event days at the urban backbround background site compared with the rural site. This is associated with a higher abundance of condensable species in the urban environment, which enhances the nucleation and growth process. This effect though is limited as particle size increases and NPF events have a greater effect on the overall $N_{<100 \text{ nm}}$ in the rural areas, compared to urban, as calculated by the NSF. The effect becomes even smaller at the kerbside as the number of background particles emitted by traffic is a lot greater.

The occurrence of NPF events at the highly polluted Marylebone Road site is at first sight surprising given the elevated condensation sink. This must be counteracted by an abundance of

condensable material, which is surprising given the generally modest rate of atmospheric oxidation 564 processes in comparison to residence times in a street canyon (Harrison, 2017). However, Giorio et 565 al. (2015), using Aerosol Time-of-Flight Mass Spectrometry, reported rapid chemical processes 566 567 within the Marylebone Road street canyon leading to production of secondary particulate matter from road traffic emissions. They postulated that this resulted from very local gas to particle 568 conversion from vehicle-emitted pollutants. Condensation of such reaction products upon pre-569 existing particles could explain the enhanced particle growth rates observed at Marylebone Road 570 (Figure 3). 571

572

573

574

575

576

577

578

Finally, particle survival probability was found to decrease moving from rural to urban areas. While formation and initial growth of new particles is increased in urban areas, their survivability reduces as their size increases. The probability of particles to survive to greater sizes was found to be increased in summer for all areas, which is also explained by the higher growth rate. The probability is also different depending upon the origin of the air masses and is related to conditions specific for each area.

579

In the present work, the effects of atmospheric conditions upon the NPF process are studied. NPF is a complex process, highly affected by meteorological conditions (local and synoptic), the chemical composition as well as the pre-existing conditions in an area. For this reason, the study of NPF events in one area cannot provide safe assumptions for other areas, as the mixture of conditions

584	found in different places is unique and alters the occurrence and development of NPF events. Thus,
585	more studies on the conditions and the trends in NPF events should be conducted to better
586	understand the effect of the numerous variables that affect those processes.
587	
588	DATA AVAILABILITY
589	Data supporting this publication are openly available from the UBIRA eData repository at
590	https://doi.org/10.25500/edata.bham.00000307.
591	
592	AUTHOR CONTRIBUTIONS
593	This study was conceived by MD and RMH who also contributed to the final manuscript. The data
594	analysis was carried out by DB with guidance from DCSB, and DB also prepared the first draft of
595	the manuscript. FDP provided advice on the analysis.
596	
597	COMPETING INTERESTS
598	The authors have no conflict of interests.
599	
600	ACKNOWLEDGEMENT
601	The authors acknowledge financial support (to DCSB) from the National Environment Research
602	Council's funding of the National Centre for Atmospheric Science (NCAS) (Grant Number
603	R8/H12/83/011).
604	31

REFERENCES

- Alam, A., Shi, J. P. and Harrison, R. M.: Observations of new particle formation in urban air,
- Journal of Geophysical Research: Atmospheres, 108, 4093-4107, doi:10.1029/2001JD001417, 607
- 2003. 608
- 609
- 610 Atkinson, R. W., Fuller, G. W., Anderson, H. R., Harrison, R. M. and Armstrong, B.: Urban
- ambient particle metrics and health: A time-series analysis, Epidemiology, 21, 501-511, 2010. 611
- 612
- Beddows, D. C. S., Harrison, R. M., Green, D. C., and Fuller, G. W.: Receptor modelling of both 613
- particle composition and size distribution from a background site in London, UK, Atmos, Chem. 614
- Phys., 15, 10107-10125, 2015. 615
- 616
- Berndt, T., Böge, O., and Stratmann, F.: Formation of atmospheric H2SO4H2O particles in the 617
- absence of organics: A laboratory study, Geophys. Res. Lett., 33, 2–6, 2006. 618
- 619 620
 - Bianchi, F., Trostl, J., Junninen, H., Frege, C., Henne, S., Hoyle, C. R., Molteni, U., Herrmann, E.,
- Adamov, A., Bukowiecki, N., Chen, X., Duplissy, J., Gysel, M., Hutterli, M., Kangasluoma, J., 621
- Kontkanen, J., Kurten, A., Manninen, H. E., Munch, S., Perakyla, O., Petaja, T., Rondo, L., 622
- Williamson, C., Weingartner, E., Curtius, J., Worsnop, D. R., Kulmala, M., Dommen, J., and 623
- Baltensperger, U.: New particle formation in the free troposphere: A question of chemistry and 624
- timing, Science, 352, 1109-1112, 2016. 625
- 626

629

- Bigi, A. and Harrison, R. M.: Analysis of the air pollution climate at a central urban background 627
- 628 site, Atmos. Environ., 44, 2004–2012, 2010.
- Brines, M., Dall'Osto, M., Beddows, D. C. S., Harrison, R. M., and Querol, X.: Simplifying aerosol 630
- size distributions modes simultaneously detected at four monitoring sites during SAPUSS, Atmos. 631
- Chem. Phys., 14, 2973-2986, 2014. 632
- 634
- Carslaw, D. C. and Ropkins, K.: openair An R package for air quality data analysis, Environ.
- Modell. Softw., 27-28, 52-61, 2012. 635
- 636

- Charron, A., Birmili, W., and Harrison, R. M.: Factors influencing new particle formation at the 637
- 638 Harwell, United Kingdom, J. Geophys. Res., Atmospheres,
- doi:10.1029/2007JD0084252007. 639
- 640
- Charron, A., Birmili, W., and Harrison, R. M.: Fingerprinting particle origins according to their size 641
- distribution at a UK rural site, J. Geophys. Res., Atmospheres, 113, D07202, 642
- doi:10.1029/2007JD008562, 2008. 643

- 644 Charron, A., Degrendele, C., Laongsri, B., and Harrison, R. M.: Receptor modelling of secondary
- and carbonaceous particulate matter at a southern UK site, Atmos. Chem. Phys., 13, 1879-1894,
- 646 2013.
- 647
- Charron, A. and Harrison, R. M.: Primary particle formation from vehicle emissions during exhaust
- dilution in the roadside atmosphere, Atmos. Environ., 37, 4109–4119, 2003.
- 650
- Dal Maso, M., Kulmala, M., Lehtinen, K. E. J., Mkelä, J. M., Aalto, P., and O'Dowd, C. D.:
- 652 Condensation and coagulation sinks and formation of nucleation mode particles in coastal and
- 653 boreal forest boundary layers, J.Geophys. Res., Atmospheres, 107, doi: 10.1029/2001JD001053,
- 654 2002.
- 655
- Dal Maso, M., Kulmala, M., Riipinen, I., Wagner, R., Hussein, T., Aalto, P. P., and Lehtinen, K. E.
- 657 J.: Formation and growth of fresh atmospheric aerosols: Eight years of aerosol size distribution data
- 658 from SMEAR II, Hyytiälä, Finland, Boreal Environ, Res., 10, 323–336, 2005.
- 659 660 Γ
 - Dall'Osto, M., Beddows, D. C. S., Asmi, A., Poulain, L., Hao, L., Freney, E., Allan, J. D.,
- Canagaratna, M., Crippa, M., Bianchi, F., de Leeuw, G., Eriksson, A., Swietlicki, E., Hansson, H.
- 662 C., Henzing, J. S., Granier, C., Zemankova, K., Laj, P., Onasch, T., Prevot, A., Putaud, J. P.,
- 663 Sellegri, K., Vidal, M., Virtanen, A., Simo, R., Worsnop, D., O'Dowd, C., Kulmala, M., and
- 664 Harrison, R. M.: Novel insights on new particle formation derived from a pan-european observing
- 665 system, Sci. Rep., 8, 1482, 2018.
- 666
- 667 Dall'Osto, M., Beddows, D. C. S., Pey, J., Rodriguez, S., Alastuey, A., M. Harrison, R., and
- 668 Querol, X.: Urban aerosol size distributions over the Mediterranean city of Barcelona, NE Spain,
- 669 Atmos. Chem. Phys., 12, 10693-10707, 2012.
- 670
- 671 Dall'Osto, M., Querol, X., Alastuey, A., O'Dowd, C., Harrison, R. M., Wenger, J., and Gómez-
- Moreno, F. J.: On the spatial distribution and evolution of ultrafine particles in Barcelona, Atmos.
- 673 Chem. Phys., 13, 741–759, 2013.
- 674
- 675 Dall'Osto, M., Thorpe, A., Beddows, D. C. S., Harrison, R. M., Barlow, J. F., Dunbar, T., Williams,
- 676 P. I., and Coe, H.: Remarkable dynamics of nanoparticles in the urban atmosphere, Atmos. Chem.
- 677 Phys., 11, 6623–6637, 2011.
- 678
- 679 Dall'Osto, M., Beddows, D. C. S., Tunved, P., Krejci, R., Ström, J., Hansson, H. C., Yoon, Y. J.,
- 680 Park, K. T., Becagli, S., Udisti, R., Onasch, T., Ódowd, C. D., Simó, R., and Harrison, R. M.: Arctic
- sea ice melt leads to atmospheric new particle formation, Sci.Rep., 7, 0–10, 2017.
- 682 683

- 684 Dameto de España, C., Wonaschütz, A., Steiner, G., Rosati, B., Demattio, A., Schuh, H. and
- 685 Hitzenberger, R.: Long-term quantitative field study of New Particle Formation (NPF) events as a
- 686 source of Cloud Condensation Nuclei (CCN) in the urban background of Vienna, Atmos. Environ.,
- 687 164, 289–298, 2017.

689 Davidson, C. I., Phalen, R. F., and Solomon, P. A.: Airborne particulate matter and human health: A review, Aerosol Sci. Technol., 39, 737–749, 2005.

691

692 Draxler, R. R., and Hess, G. D.: An Overview of the HYSPLIT_4 Modelling System for 693 Trajectories, Dispersion, and Deposition, Australian Meteorolog. Mag., 47, 295–308, 1998.

694

- 695 Ehn, M., Vuollekoski, H., Petäjä, T., Kerminen, V.-M., Vana, M., Aalto, P., de Leeuw, G.,
- 696 Ceburnis, D., Dupuy, R., O'Dowd, C. D., and Kulmala, M.: Growth rates during coastal and marine
- 697 new particle formation in western Ireland, J. Geophys. Res., 115, D18218,
- 698 http://dx.doi.org/10.1029/2010JD014292, 2010.

699

- 700 Fiedler, V., Dal Maso, M., Boy, M., Aufmhoff, H., Hoffmann, J., Schuck, T., Birmili, W., Arnold,
- 701 F., and Kulmala, M.: The contribution of sulphuric acid to atmospheric particle formation and
- growth: a comparison between boundary layers in Northern and Central Europe, Atmos. Chem.
- 703 Phys. Discuss. 5, 573-605, 2005.

704

705 FuksFuchs, N. A. and Sutugin, A. G.: Highly Dispersed Aerosols, Foreign Sci. and Technol. 706 Center, 1–86, 1971.

707

- 708 Gentner, D. R., Isaacman, G., Worton, D. R., Chan, A. W. H., Dallmann, T. R., Davis, L., Liu, S.,
- 709 Day, D. A., Russell, L. M., Wilson, K. R., Weber, R., Guha, A., Harley, R. A., and Goldstein, A.
- 710 H.: Elucidating secondary organic aerosol from diesel and gasoline vehicles through detailed
- 711 characterization of organic carbon emissions, Proc. Natl. Acad. Sci., 109, 18318–18323, 2012.

712

- 713 Hama, S. M. L., Cordell, R. L., Kos, G. P. A., Weijers, E. P., and Monks, P. S.: Sub-micron particle
- 714 number size distribution characteristics at two urban locations in Leicester, Atmos. Res., 194, 1–16.
- 715 2017.

716

- 717 Harrison, R. M., Beddows, D. C. S., Alam, M. S., Singh, A., Brean, J., and R. Xu: Interpretation of
- 718 Particle number size distributions measured across an urban area during the FASTER campaign, in
- 719 preparation, 2018.
- 720 Harrison, R. M.: Urban atmospheric chemistry: a very special case for study, npj Climate and
- 721 Atmos. Sci., 1, 5, 2017.

- Harrison, R. M., Shi, J. P., Xi, S., Khan, A., Mark, D., Kinnersley, R., and Yin, J.: Measurement of
- 725 number, mass and size distribution of particles in the atmosphere, Philos. Trans. A. Math. Phys.
- 726 Eng. Sci., 358, 2567–2580, 2000.
- 727
- Harrison, R. M. and Yin, J.: Particulate matter in the atmosphere: Which particle properties are important for its effects on health?, Sci. Tot. Environ., 249, 85–101, 2000.
- 730
- Harrison, R.M., Yin, J., Tilling, R.M., Cai, X., Seakins, P.W., Hopkins, J.R., Lansley, D.L.,
- 732 Lewis, A.C., Hunter, M.C., Heard, D.E., Carpenter, L.J., Creasey, D.C., Lee, J.D., Pilling, M.J.,
- 733 Carslaw, N, Emmerson, K.M., Redington, A., Derwent, R.G., Ryall, D., Mills G., and Penkett, S.A.,
- 734 Measurement and Modelling of Air Pollution and Atmospheric Chemistry in the UK West
- 735 Midlands Conurbation: Overview of the PUMA Consortium Project, Sci. Tot. Environ., 360, 5-25
- 736 2006.
- 737 738
 - 738 Hietikko, R., Kuuluvainen, H., Harrison, R. M., Portin, H., Timonen, H., Niemi, J. V., Ronkko, T.:
- 739 Diurnal variation of nanocluster aerosol concentrations and emission factors in a street canyon,
- 740 Atmos. Environ., 189, 98-106, 2018.
- 741
- 742 Iida, K., Stolzenburg, M. R., McMurry, P. H., and Smith, J. N.: Estimating nanoparticle growth
- 743 rates from size-dependent charged fractions: Analysis of new particle formation events in Mexico
- 744 City, J. Geophys. Res. Atmospheres, 113, D05207, doi:10.1029/2007JD009260, 2008.
- 745

- Jayaratne, R., Pushpawela, B., He, C., Li, H., Gao, J., Chai, F., and Morawska, L.: Observations of
- particles at their formation sizes in Beijing, China, Atmos. Chem. Phys., 17, 8825–8835, 2017.
- 749 Jeong, C.-H., Evans, G. J., McGuire, M. L., Chang, R. Y.-W., Abbatt, J. P. D., Zeromskiene, K.,
- 750 Mozurkewich, M., Li, S.-M., and Leaitch, W. R.: Particle formation and growth at five rural and
- 751 urban sites, Atmos. Chem. Phys., 10, 7979–7995, 2010.
- 752

- 753 Kalkavouras, P., Bossioli, E., Bezantakos, S., Bougiatioti, A., Kalivitis, N., Stavroulas, I.,
- 754 Kouvarakis, G., Protonotariou, A. P., Dandou, A., Biskos, G., Mihalopoulos, N., Nenes, A., and
- 755 Tombrou, M.: New particle formation in the southern Aegean Sea during the Etesians: Importance
- 756 for CCN production and cloud droplet number, Atmos. Chem. Phys., 17, 175–192, 2017.
- 758 Kecorius, S., Kivekäs, N., Kristensson, A., Tuch, T., Covert, D. S., Birmili, W., Lihavainen, H.,
- 759 Hyvärinen, A. P., Martinsson, J., Sporre, M. K., Swietlicki, E., Wiedensohler, A., and Ulevicius, V.:
- 760 Significant increase of aerosol number concentrations in air masses crossing a densely trafficked
- 761 sea area, Oceanologia, 58, 1–12, 2016.
- 762 763

- 764 Kelly, F. J. and Fussell, J. C.: Size, source and chemical composition as determinants of toxicity
- attributable to ambient particulate matter, Atmos. Environ., 60, 504–526, 2012.
- 766
- 767 Keuken, M. P., Moerman, M., Zandveld, P., Henzing, J. S., and Hoek, G.: Total and size-resolved
- 768 particle number and black carbon concentrations in urban areas near Schiphol airport (the
- 769 Netherlands), Atmos. Environ., 104, 132–142, 2015.

- 771 Kirkby, J. et al.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol
- 772 nucleation, Nature, 476(7361), pp. 429–435, 2011.

773

- 774 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagne, S.,
- 775 Ickes, L., Kurten, A., Kupc, Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S.,
- 776 Tsagkogeorgas, G., Wimmer, D., Amorim, A. A., Bianchi, F., Breitenlechner, M., David, A.,
- 777 Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud, W.,
- 778 Junninen, H., Kreiss, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R.,
- 779 Makhmutov, V., Mathot, S., Mikkila, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A.,
- 780 Pereira, P., Petaja, T., Schnitzhofer, R., Seinfeld, J. H., Sipila, M., Stozhkov, Y., Stratmann, F.,
- 781 Tome, A., Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E.,
- 782 Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U., and Kulmala M.:
- 783 Ion-induced nucleation of pure biogenic particles, Nature. Nature Publishing Group, 533, 521–526,
- 784 2016.

785

- 786 Korhonen, P., Kulmala, M., Laaksonen, A., Viisanen, Y., Mcgraw, R., and Seinfeld, J. H.: Ternary
- 787 nucleation of H₂SO₄, NH₃ and H₂O in the atmosphere, J. Geophys. Res., 104, D21, 26,349–26,353,
- 788 1999.

789

793

- 790 Kuang, C., McMurry, P. H., McCormick, A. V., and Eisele, F. L.: Dependence of nucleation rates
- 791 on sulfuric acid vapor concentration in diverse atmospheric locations, J. Geophys. Res.,
- 792 Atmospheres, 113, D10209, doi:10.1029/2007JD009253, 2008.
- 794 Kulmala, M., Dal Maso, M., Mäkelä, J. M., Pirjola, L., Väkevä, M., Aalto, P., Miikkulainen, P.,
- 795 Hämeri, K., and O'Dowd, C. D.: On the formation, growth and composition of nucleation mode
- particles, Tellus, Series B: Chem. Phys.Meteorol., 53, 479–490, 2001.
- 798 Kulmala, M., Kerminen, V.-M., Petäjä, T., Ding, A. J., and Wang, L.: Atmospheric gas-to-particle
- 799 conversion: why NPF events are observed in megacities?, Faraday Discuss., 271–288, 2017.

800

- Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petaja, T.,
- 802 Sipila, M., Schobesberger, S., Rantala, P., Franchin, A., Jokinen, T., Jarvinen, E., Aijala, M.,
- 803 Kangasluoma, J., Hakala, J., Aalto, P. P., Paasonen, P., Mikkila, J., Vanhanen, J., Aalto, J., Hakola,

- 804 H., Makkonen, U., Ruuskanen, T., Mauldin, R. L., Duplissy, J., Vehkamaki, H., Back, J.,
- 805 Kortelainen, A., Riipinen, I., Kurten, T., Johnston, M. V., Smith, J. N., Ehn, M., Mentel, T. F.,
- 806 Lehtinen, K. E. J., Laaksonen, A., Kerminen, V.-M., and Worsnop, D. R.: Direct Observations of
- 807 Atmospheric Aerosol Nucleation, Science, 339, 943–946, 2013.
- 808
- 809 Kulmala, M., Luoma, K., Virkkula, A., Petäjä, T., Paasonen, P., Kerminen, V. M., Nie, W., Qi, X.,
- 810 Shen, Y., Chi, X., and Ding, A.: On the mode-segregated aerosol particle number concentration
- 811 load: Contributions of primary and secondary particles in Hyytiälä and Nanjing, Boreal Environ.
- 812 Res., 21, 319–331, 2016.
- 813
- 814 Kulmala, M., Petäjä, T., Ehn, M., Thornton, J., Sipilä, M., Worsnop, D. R., and Kerminen, V.-M.:
- 815 Chemistry of Atmospheric Nucleation: On the Recent Advances on Precursor Characterization and
- 816 Atmospheric Cluster Composition in Connection with Atmospheric New Particle Formation,
- 817 Ann.Rev.Phys. Chem., 65, 21–37, 2014.
- 818
- Kulmala, M., Petäjä, T., Mönkkönen, P., Koponen, I. K., Dal Maso, M., Aalto, P. P., Lehtinen, K.
- 820 E. J., and Kerminen, V.-M.: On the growth of nucleation mode particles: source rates of
- 821 condensable vapor in polluted and clean environments, Atmos. Chem. Phys. Discuss., 4, 6943-
- 822 6966, 2005.
- 823
- 824 Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M.,
- 825 Aalto, P. P., Junninen, H., Paasonen, P., Riipinen, I., Lehtinen, K. E. J., Laaksonen, A., and
- 826 Kerminen, V. M.: Measurement of the nucleation of atmospheric aerosol particles, Nature
- 827 Protocols, 7, 1651–1667, 2012.
- 828
- 829 Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili, W.,
- 830 and McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: A review of
- 831 observations, J. Aerosol Sci., 35, 143–176, 2004.
- 832
- 833 Laaksonen, A., Kulmala, M., O'Dowd, C. D., Joutsensaari, J., Vaattovaara, P., Mikkonen, S.,
- 834 Lehtinen, K. E. J., Sogacheva, L., Dal Maso, M., Aalto, P., Petäjä, T., Sogachev, A., Yoon, Y. J.,
- 835 Lihavainen, H., Nilsson, D., Facchini, M. C., Cavalli, F., Fuzzi, S., Hoffmann, T., Arnold, F.,
- 836 Hanke, M., Sellegri, K., Umann, B., Junkermann, W., Coe, H., Allan, J. D., Alfarra, M. R.,
- 837 Worsnop, D. R., Riekkola, M. L., Hyötyläinen, T., and Viisanen, Y.: The role of VOC oxidation
- products in continental new particle formation, Atmos. Chem. Phys., 8, 657–2665, 2008.
- 839
- 840 Ma, N. and Birmili, W.: Estimating the contribution of photochemical particle formation to ultrafine
- particle number averages in an urban atmosphere, Sci. Tot. Environ., 512–513, 154–166, 2015.
- 842
- 843

- MacNee, W. and Donaldson, K.: Mechanism of lung injury caused by PM10 and ultrafine particles
- with special reference to COPD, Europ. Respirat. J., 21, 47S–51S, 2003.
- 846
- 847 Makkonen, R., Asmi, A., Kerminen, V. M., Boy, M., Arneth, A., Hari, P., and Kulmala, M.: Air
- 848 pollution control and decreasing new particle formation lead to strong climate warming. Atmos.
- 849 Chem. Phys., 12, 1515–1524, 2012.

- 851 Masiol, M., Harrison, R. M., Vu, T. V., and Beddows, D. C. S.: Sources of sub-micrometre particles
- near a major international airport, Atmos. Chem. Phys., 17, 12379–12403, 2017.

853

- 854 Merikanto, J., Spracklen, D. V., Mann, G. W., Pickering, S. J., and Carslaw, K. S.: Impact of
- nucleation on global CCN, Atmos. Chem. Phys., 9, 8601–8616, 2009.

856

- 857 Metzger, A., Verheggen, B., Dommen, J., Duplissy, J., Prevot, A. S. H., Weingartner, E., Riipinen,
- 858 I., Kulmala, M., Spracklen, D. V., Carslaw, K. S., and Baltensperger, U.: Evidence for the role of
- organics in aerosol particle formation under atmospheric conditions, Proc. Natl. Acad. Sci., 107,
- 860 6646–6651, 2010.

861

- 862 Minguillón, M. C., Brines, M., Pérez, N., Reche, C., Pandolfi, M., Fonseca, A. S., Amato, F.,
- 863 Alastuey, A., Lyasota, A., Codina, B., Lee, H. K., Eun, H. R., Ahn, K. H., and Querol, X.: New
- 864 particle formation at ground level and in the vertical column over the Barcelona area, Atmos. Res.,
- 865 164–165, 118–130, 2015.

866

- 867 Napari, I., Noppel, M., Vehkamäki, H., and Kulmala, M.: An improved model for ternary
- nucleation of sulfuric acid-ammonia-water, J. Chem. Phys., 116, 4221–4227, 2002.
- 869 Németh, Z., Rosati, B., Zíková, N., Salma, I., Bozó, L., Dameto de España, C., Schwarz, J., Ždímal,
- 871 V., and Wonaschütz, A.: Comparison of atmospheric new particle formation events in three Central
- 872 European cities, Atmos. Environ., 178, 191–197, 2018.

873

- 874 Németh, Z. and Salma, I.: Spatial extension of nucleating air masses in the Carpathian Basin,
- 875 Atmos. Chem. Phys., 14, 8841–8848, 2014.

- Nieminen, T., Kerminen, V.-M., Petäjä, T., Aalto, P. P., Arshinov, M., Asmi, E., Baltensperger, U.,
- 878 Beddows, D. C. S., Beukes, J. P., Collins, D., Ding, A., Harrison, R. M., Henzing, B., Hooda, R.,
- Hu, M., Hõrrak, U., Kivekäs, N., Komsaare, K., Krejci, R., Kristensson, A., Laakso, L., Laaksonen,
- 880 A., Leaitch, W. R., Lihavainen, H., Mihalopoulos, N., Németh, Z., Nie, W., O'dowd, C., Salma, I.,
- 881 Sellegri, K., Svenningsson, B., Swietlicki, E., Tunved, P., Ulevicius, V., Vakkari, V., Vana, M.,
- 882 Wiedensohler, A., Wu, Z., Virtanen, A., and Kulmala, M.: Global analysis of continental boundary
- layer new particle formation based on long-term measurements, Atmos. Chem. Phys. Discuss, 5194,

884 2018–304, 2018.

885

- 886 Nilsson, E. D., Paatero, J. and Boy, M.: Effects of air masses and synoptic weather on aerosol
- formation in the continental boundary layer, Tellus, Series B: Chem. Phys. Meteorol., 53, 462–478,
- 888 2001.

889 890

890 O'Dowd, C. D., Aalto, P., Hmeri, K., Kulmala, M., and Hoffmann, T.: Atmospheric particels from organic vapours, Nature, 416, 497–498, 2002.

892

- 893 O'Dowd, C., Jimenez, J. L., Bahreini, R., Flagan, R. C., Seinfeld, J. H., Hameri Kaarle, Pirjola, L.,
- 894 Kulmala, M., Gerard Jennings, S., and Hoffmann, T.: Marine aerosol formation from biogenic
- 895 iodine emissions, Nature, 417, 1–5, 2002.

896

- 897 Oberdurster, G.: Toxicology of ultrafine particles: in vivo studies, Philos. Trans. A. Math. Phys.
- 898 Eng. Sci., 358, 2719–2740, 2000.

899

- 900 Park, M., Yum, S. S., and Kim, J. H.: Characteristics of submicron aerosol number size distribution
- 901 and new particle formation events measured in Seoul, Korea, during 2004-2012, Asia-Pacific J.
- 902 Atmos. Sci., 51, 1-10, 2015.

903

- 904 Peng, Y., Dong, Y., Li, X., Liu, X., Dai, J., Chen, C., Dong, Z., Du, C., and Wang, Z.: Different
- 905 Characteristics of New Particle Formation Events at Two Suburban Sites in Northern China,
- 906 Atmosphere, 8, 58, 2017.

907

- 908 Penttinen, P., Timonen, K. L., Tiittanen, P., Mirme, A., Ruuskanen, J., and Pekkanen, J.: Number
- 909 concentration and size of particles in urban air: Effects on spirometric lung function in adult
- 910 asthmatic subjects, Environ. Health Perspect., 109, 319–323, 2001.

911

- 912 Petäjä, T., Mauldin, R. L., III, Kosciuch, E., McGrath, J., Nieminen, T., Paasonen, P., Boy, M.,
- 913 Adamov, A., Kotiaho, T., and Kulmala, M.: Sulfuric acid and OH concentrations in a boreal forest
- 914 site, Atmos. Chem. Phys., 9, 7435-7448, 2009.

915

- 916 Pikridas, M., Sciare, J., Freutel, F., Crumeyrolle, S., Von Der Weiden-Reinmüller, S. L., Borbon,
- 917 A., Schwarzenboeck, A., Merkel, M., Crippa, M., Kostenidou, E., Psichoudaki, M., Hildebrandt, L.,
- 918 Engelhart, G. J., Petäjä, T., Prévôt, A. S. H., Drewnick, F., Baltensperger, U., Wiedensohler, A.,
- 919 Kulmala, M., Beekmann, M., and Pandis, S. N.: In situ formation and spatial variability of particle
- 920 number concentration in a European megacity, Atmos. Chem. Phys., 15, 0219–10237, 2015.

- 922 Politis, M., Pilinis, C., and Lekkas, T. D.: Ultrafine particles (UFP) and health effects. Dangerous.
- 923 Like no other PM? Review and analysis, Global Nest J., 10, 439–452, 2008.

- 924 Rahman, M. M., Mazaheri, M., Clifford, S., and Morawska, L.: Estimate of main local sources to
- 925 ambient ultrafine particle number concentrations in an urban area, Atmos. Res., 194, 178-189,
- 926 2017.
- 927
- 928 Riccobono, F., Schobesberger, S., Scott, C. E., Dommen, J., Ortega, I. K., Rondo, L., Almeida, J.,
- 929 Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Downard, A., Dunne, E. M., Duplissy, J.,
- 930 Ehrhart, S., Flagan, R. C., Franchin, A., Hansel, A., Junninen, H., Kajos, M., Keskinen, H., Kupc,
- 931 A., Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Petäjä, T., Tsagkogeorgas, G.,
- 932 Vaattovaara, P., Viisanen, Y., Vrtala, A., and Wagner, P. E.: Oxidation Products of Biogenic
- 933 Atmospheric Particles, Science, 717, 17–722, 2014.
- 934
- 935 Riipinen, I., Sihto, S.-L., Kulmala, M., Arnold, F., Dal Maso, M., Birmili, W., Saarnio, K., Teinilä,
- 936 K., Kerminen, V.-M., Laaksonen, A., and Lehtinen, K. E. J.: Connections between atmospheric
- 937 sulphuric acid and new particle formation during QUEST III-IV campaigns in Heidelberg and
- 938 Hyytiälä, Atmos. Chem. Phys., 7, 1899–1914, 2007.
- 939
- P40 Robinson, A. L., Donahue, N. M., Shrivastava, M. K., Weitkamp, E. A., Sage, A. M., Grieshop, A.
- 941 P., Lane, T. E., Pierce, J. R., and Pandis, S. N.: Rethinking Organic Aerosols:, Science, 315, 1259-
- 942 1262, 2007.
- 943 944 **R**önk
 - 944 Rönkkö, T., Kuuluvainen, H., Karjalainen, P., Keskinen, J., Hillamo, R., Niemi, J. V., Pirjola, L.,
 - 945 Timonen, H. J., Saarikoski, S., Saukko, E., Järvinen, A., Silvennoinen, H., Rostedt, A., Olin, M.,
- 946 Yli-Ojanperä, J., Nousiainen, P., Kousa, A. and Dal Maso, M.: Traffic is a major source of
- atmospheric nanocluster aerosol, Proc. Natl. Acad. Sci., 114, 7549–7554, 2017.
- 948

- 949 Salma, I., Borsós, T., Németh, Z., Weidinger, T., Aalto, P., and Kulmala, M.: Comparative study of
- 950 ultrafine atmospheric aerosol within a city, Atmos. Environ., 92, 154–161, 2014.
- 952 Salma, I., Németh, Z., Kerminen, V. M., Aalto, P., Nieminen, T., Weidinger, T., Molnár, Á., Imre,
- 953 K., and Kulmala, M.: Regional effect on urban atmospheric nucleation, Atmos. Chem. Phys., 16,
- 954 8715-8728, 2016.
- 955
- 956 Salma, I., Varga, V., and Németh, Z.: Quantification of an atmospheric nucleation and growth
- 957 process as a single source of aerosol particles in a city, Atmos. Chem. Phys., 17, 15007-15017,
- 958 2017.

959

- 960 Samoli, E., Atkinson, R. W., Analitis, A., Fuller, G. W., Beddows, D., Green, D. C., Mudway, I. S.,
- 961 Harrison, R. M., Anderson, H. R., and Kelly, F. J.: Differential health effects of short-term exposure
- 962 to source-specific particles in London, U.K., Environ. Intl., 97, 246–253, 2016.

- 964 Seinfeld, J. H. and Pandis, S. N.: Atmospheric Chemistry and Physics: From Air Pollution to Climate Change, 3rd Ed. New Jersey, Canada, John Wiley & Sons, Inc, 2012.
- 968 Shen, X., Sun, J., Kivekäs, N., Kristensson, A., Zhang, X., Zhang, Y., Zhang, L., Fan, R., Qi, X.,
- 969 Ma, Q. and Zhou, H.: Spatial distribution and occurrence probability of regional new particle
- 970 formation events in eastern China, Atmos. Chem. Phys, 185194, pp. 587–599, 2018.
- 972 Shi, J. P., Evans, D. E., Khan, A. A., and Harrison, R. M.: Sources and concentration of
- 973 nanoparticles (<10nm diameter) in the urban atmosphere, Atmos. Environ., 35, 1193–1202, 2001.
- 975 Shi, J. P. and Harrison, R. M.: Investigation of ultrafine particle formation during diesel exhaust dilution, Environ. Sci. Technol., 33, 3730–3736, 1999.
- 978 Spracklen, D. V., Carslaw, K. S., Kulmala, M., Kerminen, V. M., Sihto, S. L., Riipinen, I.,
- 979 Merikanto, J., Mann, G. W., Chipperfield, M. P., Wiedensohler, A., Birmili, W., and Lihavainen,
- 980 H.: Contribution of particle formation to global cloud condensation nuclei concentrations, Geophys.
- 981 Res. Lett., 35, 1-5, 2008.

971

974

977

982

990

- 983 Spracklen, D. V., Carslaw, K. S., Merikanto, J., Mann, G. W., Reddington, C. L., Pickering, S.,
- 984 Ogren, J. A., Andrews, E., Baltensperger, U., Weingartner, E., Boy, M., Kulmala, M., Laakso, L.,
- 985 Lihavainen, H., Kivekäs, N., Komppula, M., Mihalopoulos, N., Kouvarakis, G., Jennings, S. G.,
- 986 O'Dowd, C., Birmili, W., Wiedensohler, A., Weller, R., Gras, J., Laj, P., Sellegri, K., Bonn, B.,
- 987 Krejci, R., Laaksonen, A., Hamed, A., Minikin, A., Harrison, R. M., Talbot, R., and Sun, J.:
- 988 Explaining global surface aerosol number concentrations in terms of primary emissions and particle
- 989 formation, Atmos.Chem.Phys., 10, 4775–4793, 2010.
- 991 Sutton, M. a.A., Place, C. J., Eager, M., Fowler, D., and Smith, R. I.: Assessment of the magnitude
- of ammonia emissions in the UK, Atmos. Environ., 29, 1393–1411, 1995.
- 994 Tröstl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege, C.,
- 995 Bianchi, F., Wagner, R., Simon, M., Lehtipalo, K., Williamson, C., Craven, J. S., Duplissy, J.,
- 996 Adamov, A., Almeida, J., Bernhammer, A.-K., Breitenlechner, M., Brilke, S., Dias, A., Ehrhart, S.,
- 997 Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Gysel, M., Hansel, A., Hoyle, C. R., Jokinen, T.,
- 998 Junninen, H., Kangasluoma, J., Keskinen, H., Kim, J., Krapf, M., Kürten, A., Laaksonen, A.,
- 999 Lawler, M., Leiminger, M., Mathot, S., Möhler, O., Nieminen, T., Onnela, A., Petäjä, T., Piel, F. M.,
- 1000 Miettinen, P., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Sengupta, K., Sipilä, M.,
- 1001 Smith, J. N., Steiner, G., Tomè, A., Virtanen, A., Wagner, A. C., Weingartner, E., Wimmer, D.,
- 1002 Winkler, P. M., Ye, P., Carslaw, K. S., Curtius, J., Dommen, J., Kirkby, J., Kulmala, M., Riipinen,

- 1003 I., Worsnop, D. R., Donahue, N. M., and Baltensperger, U.: The role of low-volatility organic
- 1004 compounds in initial particle growth in the atmosphere, Nature, 533, 527–531, 2016.
- 1005
- 1006 Von Bismarck-Osten, C., Birmili, W., Ketzel, M., Massling, A., Petäjä, T., and Weber, S.:
- 1007 Characterization of parameters influencing the spatio-temporal variability of urban particle number
- size distributions in four European cities, Atmos. Environ., 77, 415–429, 2013.
- 1009
- 1010 Wang, Z., Wu, Z., Yue, D., Shang, D., Guo, S., Sun, J., Ding, A., Wang, L., Jiang, J., Guo, H., Gao,
- 1011 J., Cheung, H. C., Morawska, L., Keywood, M., and Hu, M.: New particle formation in China:
- 1012 Current knowledge and further directions, Sci. Tot. Environ., 258–266, 2017.
- 1013
- 1014 Wehner, B., Siebert, H., Stratmann, F., Tuch, T., Wiedensohler, A., Petäjä, T., Dal Maso, M., and
- 1015 Kulmala, M.: Horizontal homogeneity and vertical extent of new particle formation events, Tellus,
- 1016 Series B: Chem. Phys. Meteorol., 59, 362–371, 2007.
- 1017

- 1018 Woo, K. S., Chen, D. R., Pui, D. Y. H., and McMurry, P. H.: Measurement of Atlanta aerosol size
- 1019 distributions: Observations of lutrafine particle events, Aerosol Sci. Technol., 34, 5–87, 2001.
- 1021 Xiao, S., Wang, M. Y., Yao, L., Kulmala, M., Zhou, B., Yang, X., Chen, J. M., Wang, D. F., Fu, O.
- 1022 Y., Worsnop, D. R., and Wang, L.: Strong atmospheric new particle formation in winter in urban
- 1023 Shanghai, China, Atmos. Chem. Phys., 15, 1769–1781, 2015.
- 1024
- 1025 Yli-Juuti, T., Nieminen, T., Hirsikko, A., Aalto, P. P., Asmi, E., Hõrrak, U., Manninen, H. E.,
- 1026 Patokoski, J., Dal Maso, M., Petäjä, T., Rinne, J., Kulmala, M., and Riipinen, I.: Growth rates of
- 1027 nucleation mode particles in Hyytiälä during 2003-2009: Variation with particle size, season, data
- analysis method and ambient conditions, Atmos. Chem. Phys., 11, 12865–12886, 2011.
- 1029
- 1030 Yue, D. L., Hu, M., Zhang, R. Y., Wang, Z. B., Zheng, J., Wu, Z. J., Wiedensohler, A., He, L. Y.,
- 1031 Huang, X. F., and Zhu, T.: The roles of sulfuric acid in new particle formation and growth in the
- 1032 mega-city of Beijing, Atmos. Chem. Phys., 10, 4953–4960, 2010.
- 1033
- 1034 Zhang, X., Zhang, Y., Sun, J., Zheng, X., Li, G., and Deng, Z.: Characterization of particle number
- 1035 size distribution and new particle formation in an urban environment in Lanzhou, China, J. Aerosol
- 1036 Sci., 103, 53-66, 2017.
- 1037 1038
- 1039
- 1040 1041
- 1042
- 1043

1044	TABLE LI	EGENDS:
1045		
1046	Table 1:	Number of NPF events per site.
1047		
1048	Table 2:	Annual and seasonal NSF for all areas of study.
1049		
1050		
1051		
1052	FIGURE I	ÆGENDS:
1053		
1054	Figure 1:	Map of the measuring stations.
1055	Figure 2:	Number of NPF events per season for all seven years of the present study (Winter –
1056	rigure 2.	DJF; Spring – MAM; Summer – JJA; Autumn – SON) at Harwell (rural), N.
1050		Kensington (urban background) and Marylebone Road (urban roadside).
1057		Kensington (urban background) and war yiebone Road (urban roadside).
1058	Figure 3:	Growth rate per season at the three sites.
1059	rigure 3.	Growth rate per season at the three sites.
1061	Figure 4:	Diurnal variation of N _{16-20nm} at each site: annual average and NPF event days.
1062	riguic 4.	Diditial variation of 1416-20mm at each site. Aimidal average and 1411 event days.
1063	Figure 5:	Map and frequency of incoming air mass origin – average and for NPF events per site.
1064	rigure 3.	wap and requerey of meoning an mass origin average and for 1411 events per site.
1065	Figure 6:	Growth rate per incoming air mass at each of the sites.
1066	riguite 0.	Growth rate per incoming an mass at each of the sites.
1067	Figure 7:	Survival parameter P (a) per season, (b) for regional and local events (for Marylebone
1067	rigure /:	Road) is regional for all 3 sites and (c) by incoming air mass origin.
		Road) is regional for all 3 sites and (c) by incoming air mass origin.
1069		

 Table 1: Number of NPF events per site.

	Harwell	N.	Marylebone	Regional	Regional
		Kensington	Road	(Background sites)*	(All 3 sites)**
2009	9	0	4	0	0
2010	29	22	22	11	9
2011	15	10	23	4	1
2012	8	28	12	3	0
2013	25	23	27	13	11
2014	29	34	13	18	6
2015	25	22	18	11	10
Overall	140	139	119	60	37

^{*} Refers to events occurring simultaneously at Harwell and N. Kensington
** Refers to events which occur simultaneously at all three sites

Table 2: Annual and seasonal NSF for all areas of study.

	Harwell	N. Kensington	Marylebone Road
NSF _{NUC} (Spring)	2.04	2.03	1.2 <u>0</u>
NSF _{NUC} (Summer)	2.01	1.72	1.26
NSF _{NUC} (Year)	2.25	1.86	1.26
NSF _{GEN} (Spring)	1.1 <mark>0</mark>	1.07	1.02
NSF _{GEN} (Summer)	1.18	1.11	1.01
NSF _{GEN} (Year)	1.10	1.06	1.02

Commented [DB(wISAS+19]: Addressing comment 6 (1st)

Commented [DB(wISAS+20]: Addressing comment 6 (1st)

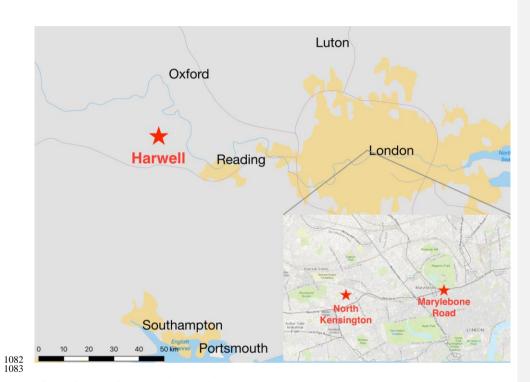


Figure 1: Map of the measuring stations.

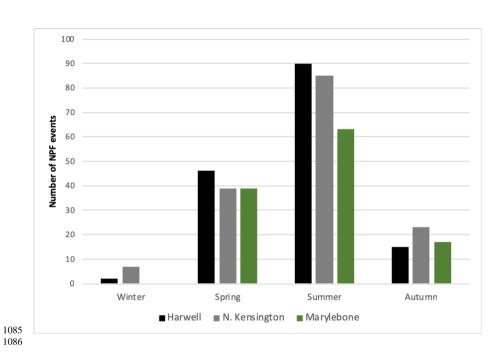


Figure 2: Number of NPF events per season for all seven years of the present study (Winter – DJF; Spring – MAM; Summer – JJA; Autumn – SON) at Harwell (rural), N.Kensington (urban background) and Marylebone Road (urban roadside).

 Commented [DB(wISAS+21]: Addressing comment 8 (1st)

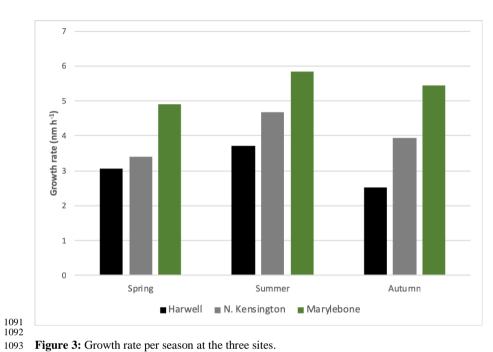


Figure 3: Growth rate per season at the three sites.

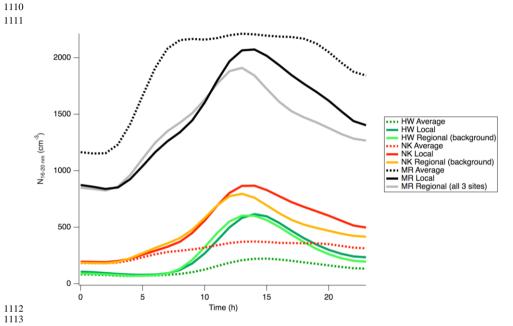


Figure 4: Diurnal variation of $N_{16\text{-}20\text{nm}}$ at each site: annual average and NPF event days.

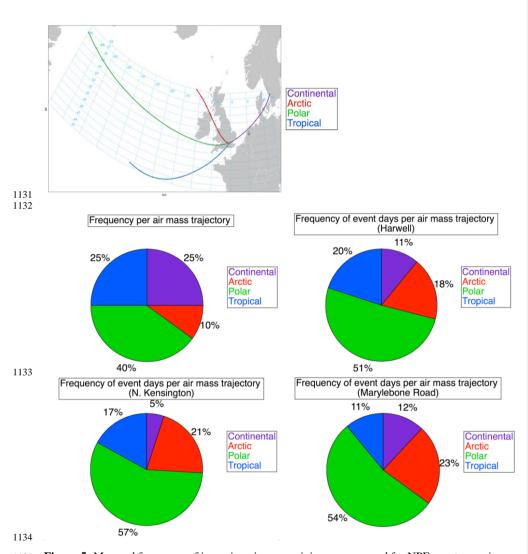


Figure 5: Map and frequency of incoming air mass origin – average and for NPF events per site.



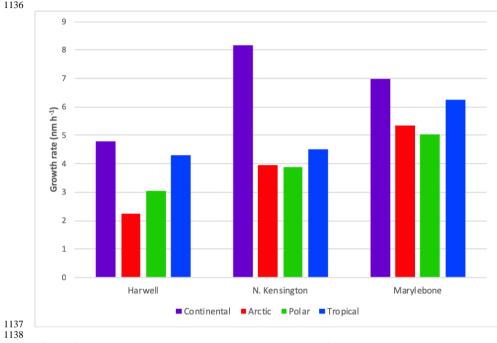
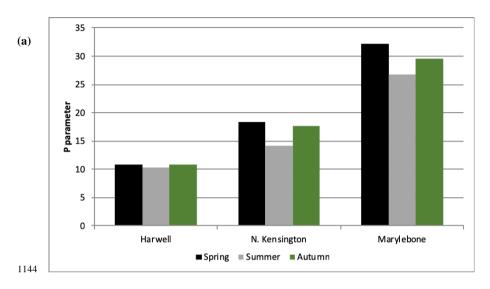
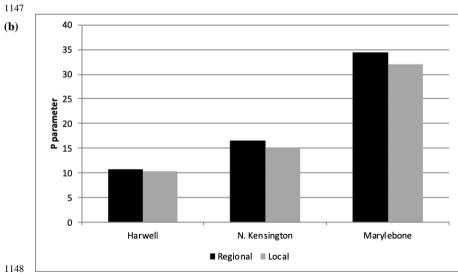


Figure 6: Growth rate per incoming air mass origin at each of the sites.





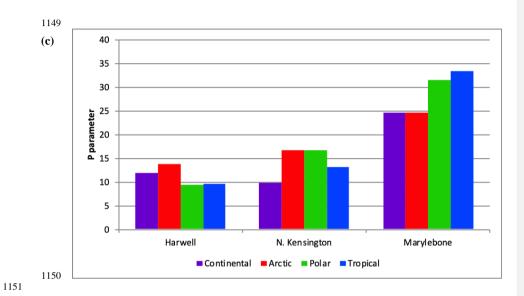


Figure 7: Survival parameter P (a) per season, (b) for regional and local events (for Marylebone Road regional is for all 3 sites) and (c) by incoming air mass origin.

SUPPLEMENTARY INFORMATION

- 3 Analysis of New Particle Formation (NPF) Events at Nearby
- 4 Rural, Urban Background and Urban Roadside Sites
- 6 Dimitrios Bousiotis, Manuel Dall'Osto, David C.S. Beddows and
- 7 Roy M. Harrison

2

Table S1: Data availability per season (all numbers are percentages of available data)

	Harwell			N. Kensington				Marylebone Road				
	Winter	Spring	Summer	<u>Autumn</u>	Winter	Spring	Summer	<u>Autumn</u>	Winter	Spring	Summer	<u>Autumn</u>
2009	<u>15</u>	97	<u>10</u>	80	<u>57</u>	97	100	100	100	<u>65</u>	86	<u>68</u>
2010	<u>37</u>	<u>53</u>	<u>100</u>	<u>95</u>	<u>58</u>	<u>87</u>	<u>93</u>	100	<u>46</u>	100	<u>87</u>	<u>86</u>
2011	<u>72</u>	<u>75</u>	<u>99</u>	<u>73</u>	<u>89</u>	<u>87</u>	<u>73</u>	<u>89</u>	<u>79</u>	99	<u>100</u>	<u>67</u>
2012	<u>82</u>	86	100	<u>95</u>	<u>56</u>	88	99	86	0	0	87	<u>66</u>
2013	<u>91</u>	<u>70</u>	99	100	84	92	98	98	<u>57</u>	92	84	100
2014	<u>97</u>	<u>62</u>	<u>99</u>	<u>99</u>	<u>84</u>	<u>78</u>	<u>97</u>	<u>98</u>	<u>89</u>	<u>79</u>	<u>76</u>	99
2015	<u>77</u>	100	<u>61</u>	<u>70</u>	80	99	<u>65</u>	100	74	100	<u>98</u>	100

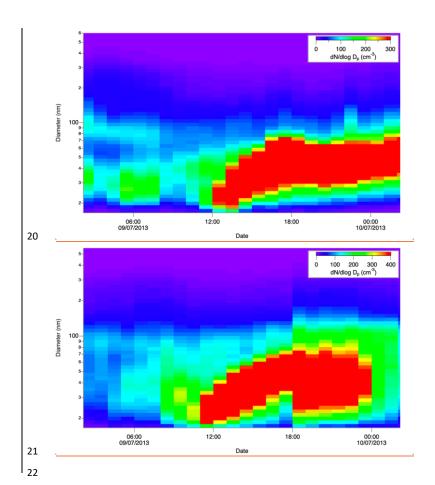
Commented [DB(wISAS+1]: Addressing comment 4 (1st)

Table S21: Conditions per air mass origin for NPF event days (April – October average in parenthesis) for all areas of study.

		Harwell		
	Continental	Arctic	Polar	Tropical
Condensation sink	5.05E-03 (5E-	2.71E-03 (3.32E-	2.57E-03	3.19E-03
(s ⁻¹)	03)	03)	(2.87E-03)	(2.87E-03)
Wind speed (m s ⁻¹)	3.52 (3.63)	3.87 (3.47)	3.64 (3.69)	3.74 (4.17)
Temperature (°C)	15.5 (13.6)	12.2 (11.5)	13.6 (13.1)	16.3 (15)
SO_2 (µg m ⁻³)	1.87 (1.81)	1.11 (1.82)	1.11 (1.27)	1.22 (1.36)
NO_x (µg m ⁻³)	9.58 (13.9)	5.49 (8.01)	4.66 (7.2)	5.81 (7.69)
SO_4^{2-} (µg m ⁻³)	2.70 (3.3)	1.37 (2.05)	1.44 (1.64)	1.37 (1.57)
Particulate OC (μg m ⁻³)	2.85 (2.88)	1.35 (1.59)	1.52 (1.63)	1.98 (1.76)

	N	orth Kensington		
	Continental	Arctic	Polar	Tropical
Condensation sink	7.20E-03	5.20E-03 (6.37E-	5.40E-03	4.89E-03
(S ⁻¹)	(9.35E-03)	03)	(6.38E-03)	(6.32E-03)
Wind speed (m s ⁻¹)	3.89 (3.44)	3.92 (3.65)	4.46 (4.2)	4.74 (4.44)
Temperature (°C)	18.4 (15)	12.7 (13.1)	15.5 (14.6)	17 (16.4)
SO_2 (µg m ⁻³)	1.68 (2.23)	1.33 (1.89)	1.73 (1.75)	1.74 (1.72)
NO_x (µg m ⁻³)	33.5 (55)	28.5 (39.2)	30.3 (39.4)	24 (34.9)
SO ₄ ²⁻ (μg m ⁻³)	1.93 (2.23)	0.95 (1.36)	0.98 (1.13)	1.30 (1.47)
Particulate OC (μg	3.84 (4.90)	2.24 (2.95)	2.81 (2.96)	2.43 (3.03)
m ⁻³)				

		Marylebone		
	Continental	Arctic	Polar	Tropical
Condensation sink	1.65E-02	1.16E-02 (1.57E-	1.4E-02 (2.14E-	1.82E-02
(S ⁻¹)	(1.96E-02)	02)	02)	(2.39E-02)
Wind speed (m s ⁻¹)	3.92 (3.41)	3.50 (3.64)	3.84 (4.13)	4.77 (4.4)
Temperature (°C)	17.4 (15.2)	13.4 (13.4)	15.3 (14.8)	16.9 (16.3)
SO_2 (µg m ⁻³)	4.99 (6.39)	4.31 (5.63)	5.38 (7.43)	6.95 (8.17)
NO_x (µg m ⁻³)	172 (250)	139 (214)	191 (303)	269 (336)
SO ₄ ²⁻ (μg m ⁻³)	3.24 (3.35)	1.47 (1.6)	1.52 (1.61)	1.24 (1.8)
Particulate OC (μg m ⁻³)	6.03 (6.91)	3.81 (4.73)	4.67 (5.97)	5.31 (6.6)



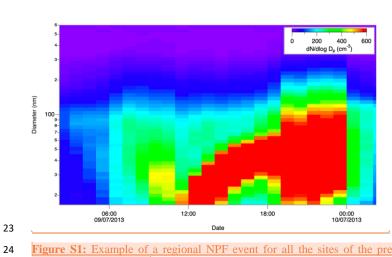
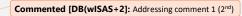
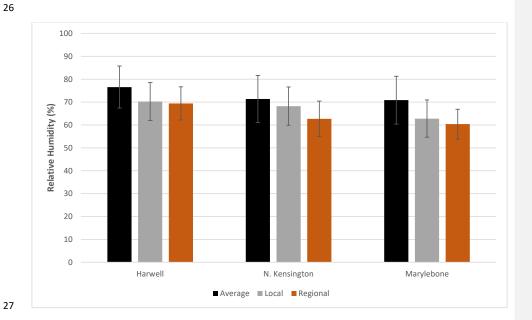
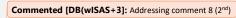
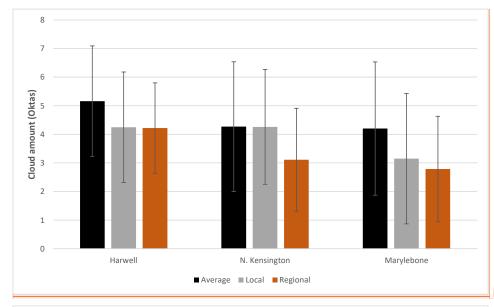


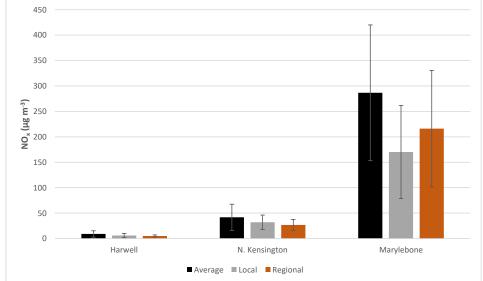
Figure S1: Example of a regional NPF event for all the sites of the present study (from top to bottom is HW, NK, MR)

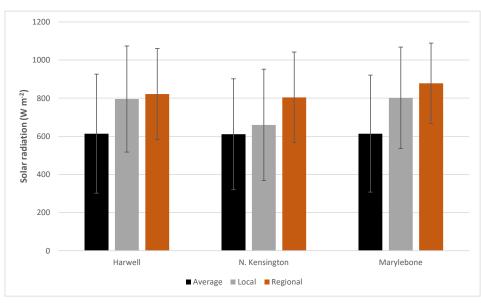


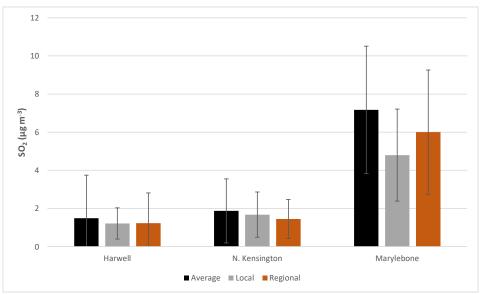


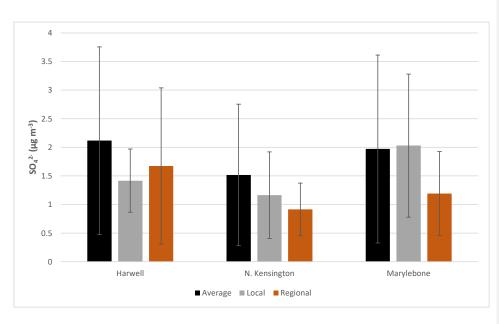


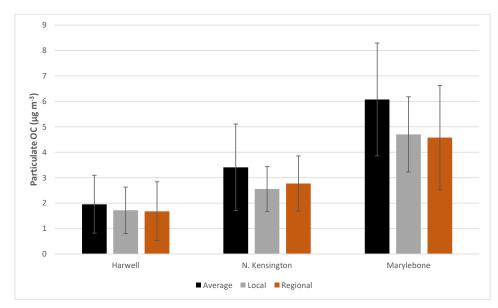












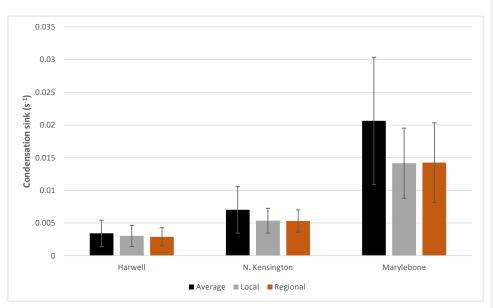
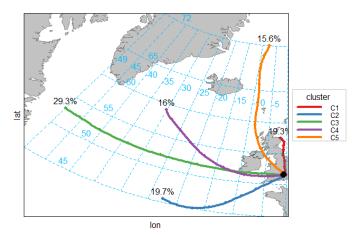
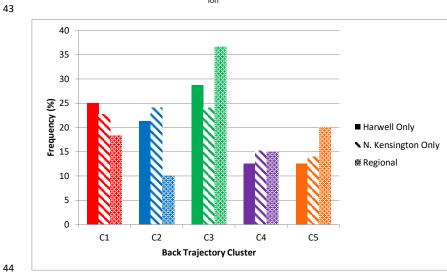
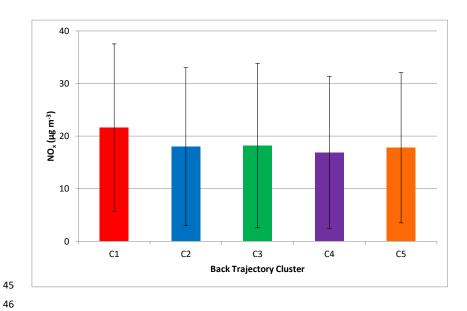
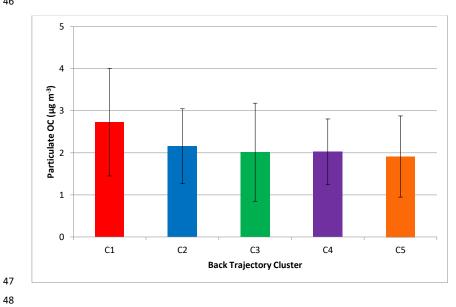


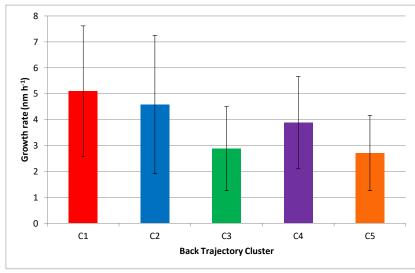
Figure S24: Average and NPF event conditions for Harwell and N. Kensington and Marylebone Road. On Marylebone Road, Regional events' conditions refer to the Regional event days for all 3 sites.

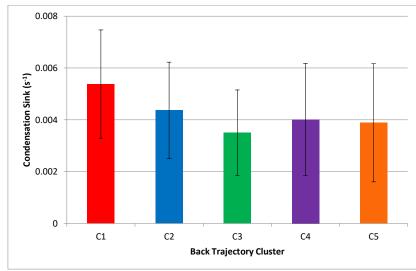


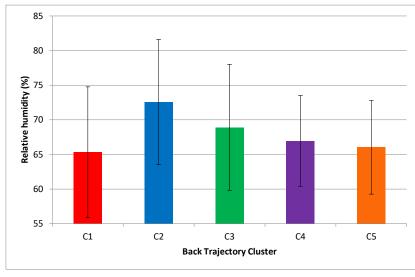


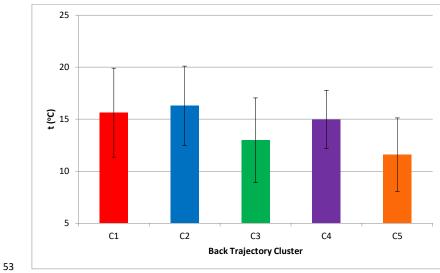


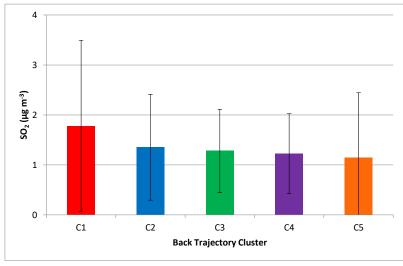












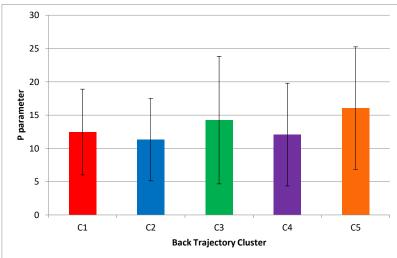


Figure S32: Air mass origin frequency and conditions for NPF events in Harwell and N. Kensington.

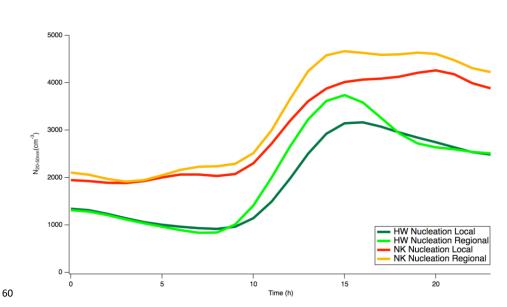


Figure S3: Diurnal variation of $N_{20-50nm}$ for Harwell and N. Kensington during local and regional NPF events.

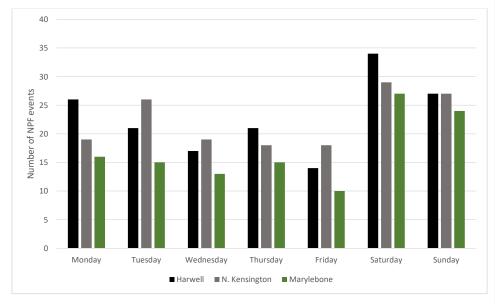
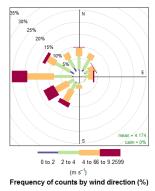


Figure S44: Weekly variation of NPF events in Marylebone Road.



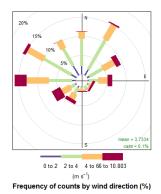


Figure So: Wind profile for Local (left) and Regional (right) NPF events in Marylebone Road.