Tropospheric HONO Distribution and Chemistry in the Southeast U.S.

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Abstract

Here we report the measurement results of nitrous acid (HONO) and a suite of relevant parameters on the NCAR C-130 research aircraft in the Southeast U.S. during the NOMADSS 2013 summer field study. The daytime HONO concentration ranged from low parts per trillion by volume (pptv) in the free troposphere (FT) to mostly within 5 - 15 pptv in the background planetary boundary layer (PBL). There was no discernible vertical HONO gradient above the lower flight altitude of 300 m in the PBL, and the transport of ground surface HONO was found not a significant contributor to the tropospheric HONO budget. The total *in situ* HONO source, mean (± 1 SD) was calculated 53 (± 21) pptv h^{-1} during the day. The upper limit contribution from NO_x-related reactions was 10 (± 5) pptv h^{-1}, and the contribution from photolysis of particulate nitrate (pNO_3) was 38 (± 23) pptv h^{-1}, based on the measured pNO_3 concentrations and the median pNO_3 photolysis rate constant of 2.0×10^{-4} s^{-1} determined in laboratory using ambient aerosol samples. The photolysis of HONO contributed to less than 10% of the primary OH source. However, a recycling NO_x source via pNO_3 photolysis was equivalent to ~ 2.3×10^{-6} mol m^{-2} h^{-1} in the air column within the PBL, a considerable supplementary NO_x source in the low-NO_x background area. Up to several tens pptv of HONO were observed in power plant and urban plumes during the day, mostly produced *in situ* from precursors including NO_x and pNO_3. Finally, there was no observable accumulation of HONO in the nocturnal residual layer and the nocturnal FT in the background Southeast U.S., with an increase in HONO/NO_x ratio of $\leq 3 \times 10^{-4}$ hr^{-1} after sunset.

1 Introduction

Extensive field studies at ground sites have shown that gas-phase nitrous acid (HONO) exists at much higher levels than expected during the day, with a mixing ratio of HONO up to several parts per billion by volume (ppbv) in the urban atmosphere (Acker et al., 2006; Villena et al., 2011) and up to several hundred parts per trillion by volume (pptv) in rural environments (Acker et al., 2006; Kleffmann et al., 2003; Zhang et al., 2009; Zhou et al., 2002, 2011). At the observed concentrations, HONO photolysis (R1) becomes an important or even a major OH primary source in both urban (Elshorbany et al., 2010; Villena et al., 2011) and rural environments near the ground surface (Acker et al., 2006; He et al., 2006; Kleffmann et al., 2003; Zhou et al., 2002, 2011).

$$HONO + h\nu \rightleftharpoons OH + NO \quad \text{(R1, R-1)}.$$

The OH radical is responsible for the removal of primary pollutants, and plays a crucial role in the formation of secondary pollutants, such as O_3 and aerosol (Finlayson-Pitts and Pitts, 2000), and thus HONO plays an important role in atmospheric chemistry.

The removal processes of HONO from the troposphere are relatively well understood, including mainly photolysis, reaction with the OH radical and surface deposition. Photolysis is the dominant sink for HONO during the day (Kleffmann et al., 2003; Oswald et al., 2015; Zhang et al., 2009, 2012), and dry deposition is the major HONO loss pathway at night, especially over wet surfaces (He et al., 2006;VandenBoer et al., 2015). However, HONO sources in the planetary boundary layer (PBL) are numerous. HONO is directly emitted from combustion processes, such as automobile emissions (Li et al., 2008b) and biomass burning (Burling et al., 2010; Trentmann et al., 2003). Soil emission via nitrification or denitrification is another source of HONO, which might be important in agriculture region (Maljanen et al., 2013; Oswald et al., 2013; Su et al., 2011). Due to the relatively short photolytic lifetime of HONO, in the order of 10 min around summer noontime, the impacts of the direct emission on HONO distribution and chemistry is highly localized and limited to the source region.

HONO is a unique species that is produced through heterogeneous reactions of different precursors, such as NO_2 and HNO_3, on surfaces (R2 - R3):

$$NO_2 + H_2O \text{ (or organics)} \xrightarrow{\text{surface}} HONO \quad \text{(R2)}$$

$$HNO_3(s) + h\nu \xrightarrow{\text{organics},H_2O} HONO + NO_2 \quad \text{(R3)}$$

Heterogeneous reactions of NO_2 with organics (R2) on the surfaces have been found to be greatly accelerated by sunlight through photosensitization (George et al., 2005; Kleffmann,

2007; Stemmler et al., 2006, 2007) and these reactions on ground surfaces are likely the major daytime HONO source in urban environments (Acker et al., 2006; Villena et al., 2011; Wong et al., 2011). Laboratory studies have confirmed that HNO_3 undergoes photolysis in sunlight at rates 2 - 3 orders of magnitude greater on the surface than in the gas phase (Baergen and Donaldson, 2013; Du and Zhu, 2011; Ye et al., 2016a, b; Zhou et al., 2003; Zhu et al., 2008), producing NO_x and HONO. In low-NO_x environments, photolysis of nitric acid/nitrate deposited on the surface has been proposed to be the major daytime HONO source near the ground surface (Ye et al., 2016b; Zhou et al., 2003, 2011).

Several processes within an air mass may lead to volume, or *in situ*, production of HONO. The OH+NO reaction (R-1) in the gas phase may be a significant HONO source in high NO_x and photochemically reactive atmospheres (Kleffmann, 2007; Neuman et al., 2016; Villena et al., 2011), but becomes negligible in low-NO_x environments(Li et al., 2014; Ye et al., 2016b). Two additional gas-phase reactions have been proposed to produce HONO within the air column: between excited NO_2 (NO_2^*) and water vapor (R4) (Li et al., 2008a), and between NO_2 and the hydroperoxyl-water complex ($HO_2 \cdot H_2O$) (R5) (Li et al., 2014):

$$NO_2^* + H_2O \rightarrow HONO + OH \quad \text{(R4)}$$

$$HO_2 \cdot H_2O + NO_2 \xrightarrow{\alpha} HONO + O_2 + H_2O \quad \text{(R5)}$$

However, further laboratory evidence suggests that reaction (R4) is too slow to be important (Carr et al., 2009; Wong et al., 2011). And recent airborne observations have demonstrated that the HONO yield (α) from reaction (R5) is less than 0.03 (Ye et al., 2015). Heterogeneous reactions of NO_2 (R2) on aerosol surfaces and photolysis of particulate nitrate (pNO_3) associated with aerosol particles similar to (R3) also contribute *in situ* HONO production in the air column.

Most HONO measurements have been made at ground stations. The observed HONO concentrations reported in the literature represent the HONO levels in the lower PBL under the significant but varying influence of ground surface processes. Thus, it is difficult to distinguish the ground surface HONO sources from the *in situ* HONO sources. Measurements of vertical HONO concentration profiles and HONO fluxes have suggested that ground surfaces can be major HONO sources for the overlying atmosphere in many cases (He et al., 2006; Kleffmann et al., 2003 Stutz et al., 2002; Zhou et al., 2011; VandenBoer et al., 2013; Young et al., 2012), but not in some other cases (Villena et al., 2011). A recent HONO flux measurement has suggested that the HONO source from the forest canopy contributed ~ 60% of the measured HONO budget at the measurement height of 11 m above the forest canopy,

and the *in situ* HONO production contributed the remaining ~ 40% (Zhou et al., 2011). Similarly, observational and modeling studies implied a presence of a volume HONO source at 130-m altitude above Houston, TX (Wong et al., 2012, 2013). The relative importance of *in situ* HONO production would be expected to increase with altitude due to decreasing influence of the ground surface, at least during the day. Airborne measurements in the air mass above the altitude influenced directly by ground HONO sources should provide more direct and quantitative evidence for *in situ* HONO production in the troposphere. Indeed, the limited number of airborne measurements available have shown that HONO exists not only in substantial amounts in combustion and urban plumes (Neuman et al., 2016) but also throughout the troposphere (Li et al., 2014; Ye et al., 2015; Zhang et al., 2009).

Here we report airborne HONO measurement results and findings from five research flights in the Southeast U.S. during the NOMADSS (Nitrogen, Oxidants, Mercury and Aerosol Distributions, Sources and Sinks) 2013 summer field campaign aboard the NSF/NCAR C-130 research aircraft.

2 Experimental

NOMADSS was an airborne field study under the "umbrella" of SAS (Southeast Atmosphere Study). It consisted of nineteen research flights on board the NSF/NCAR C-130 aircraft from June 1, 2013 to July 15, 2013. Parameters observed included HONO, HNO_3, particulate nitrate, NO_x, O_3, BrO, OH radicals, HO_2 radicals, RO_2 radicals, aerosol surface area densities (size <1 μm), VOCs, photolysis frequencies, and other meteorology parameters. Table 1 summarizes the instrumentation, time resolution, detection limit, accuracy, and references for the measurements. The results from five out of nineteen flights are presented here to discuss vertical HONO distribution and HONO chemistry in the Southeast U.S. The flight tracks are shown in Figure 1.

2.1 LPAP measurements of HONO and pNO_3

HONO was measured by two long-path absorption photometric (LPAP) systems based on the Griess-Saltzman reaction (Zhang et al., 2012; Ye et al., 2016b). Briefly, ambient air was first brought into the aircraft through a heated PFA inlet, with a residence time of 0.14 s. HONO was scrubbed using de-ionized (DI) water in two 10-turn glass coil samplers and the collected nitrite was then derivatized with 5 mM sulfanilamide (SA) and 0.5 mM N-(1-Naphthyl)-ethylene-diamine (NED) in 40 mM HCl, to form an azo dye. The azo dye was then

detected by light absorption at 540 nm using an 4-channel optic fiber spectrometer (LEDSpec, WPI) with two 1-m liquid waveguide capillary flow cells (WPI). Each LPAP system ran a 30-min measurement and zero cycle, with 20 min sampling ambient air and 10 min sampling "zero-HONO" air for baseline correction, and with a 15-min time offset between the two sampling cycles. The combination of the data from the two systems provides continues HONO concentration measurement. The "zero-HONO" air was generated by directing the sample stream through a Na_2CO_3-coated denuder to remove HONO while allowing most of interfering species (NO_x, PAN, and particulate nitrite) to pass through . The absorbance signals were sampled at a rate of 1 Hz, and were averaged into 1-min or 3-min data for analysis.

Interference from NO_x, PAN, and particulate nitrite was corrected by subtracting the baseline from the ambient air signal. Due to the low collecting efficiency of these interfering species in the sampling coil and their low concentrations, the combined interference was estimated to be less than 10% of the total signal. Potential interference from peroxynitric acid (HO_2NO_2) was suppressed by heating the PFA sampling line to 50 °C. The HO_2NO_2 steady state concentration ($[HO_2NO_2]_{SS}$) was estimated to be less than 1 pptv at temperatures of 20 - 30 °C in the background PBL (Gierczak et al., 2005), and thus interference from HO_2NO_2 was negligible. Whereas in power plant and urban plumes in the PBL or biomass burning plumes in the upper free troposphere (FT), HO_2NO_2 interference was not negligible and thus a correction for HONO measurement was made. An upper-limit HO_2NO_2 response efficiency was estimated to be 0.2 for our HONO measurement systems. The estimation was made from the lowest ratio of the measured HONO to the corresponding $[HO_2NO_2]_{SS}$ in cold air masses at high altitude, assuming no HONO existed. HONO concentration were then corrected by subtracting a term of "$0.2 \times [HO_2NO_2]_{SS}$". The correction was below 10% of the measured HONO concentrations in the PBL plumes. However, there may be over-corrections in the cold free troposphere.

The lower detection limit of LPAP-HONO measurement was estimated to be $\leq$1 pptv, based on 3 times the standard deviation of the zero air signal (N >10). An overall uncertainty of $\pm(1 + 0.2\ [HONO])$ pptv was estimated, combining the uncertainties in signal acquisition and processing, air and liquid flow rates, standard preparation, and baseline correction. The accuracy of HONO measurements was confirmed by comparison with a limb-scanning differential optical absorption spectroscopy (DOAS) (Platt and Stutz, 2008) during the NOMADSS 2013 summer field study onboard the C-130 aircraft (Ye et al., 2016b). When

152 measuring in wide power plant plumes where HONO mixing ratios exceeded the lower
153 detection limits of both instruments, the agreement between these two instruments was very
154 good, within the assessed uncertainties (Ye et al., 2016b).
155 Particulate nitrate (pNO_3) was quantitatively collected with a frit disc sampler after a
156 NaCl-coated denuder to remove HNO_3 (Huang et al., 2002). The collected nitrate was reduced
157 to nitrite by a Cd column, and determined using a LPAP system (Zhang et al., 2012). "Zero-
158 pNO_3" air was generated to establish measurement baselines for pNO_3 by passing the ambient
159 air through a Teflon filter to remove aerosol particles and then a NaCl-coated denuder to
160 remove HNO_3 before reaching the sampling unit of LPAP. Potential interferences from
161 HONO, NO_x and PAN were corrected by subtracting the baselines from the ambient air
162 signals. The lower detection limit of pNO_3 was estimated to be 2 pptv, based on 3 times the
163 standard deviation of the zero-air signal (N >10). An overall uncertainty of $\pm(2 + 0.3\ [pNO_3])$
164 pptv was estimated.
165 Noisy baselines were observed when the C-130 was flying in the clouds, due to the
166 sampling of cloud droplets by our sampling systems. Because of the lack of a valid way to
167 correct for this interference, all in-cloud measurement data of HONO and pNO_3 have been
168 excluded from the data analysis. In addition, large baseline shifts were observed sometimes
169 when the flow state of the scrubbing solution was disturbed by rapid pressure changes during
170 aircraft's rapid ascending to or descending from high altitudes. The data were excluded from
171 analysis if the baseline shifts caused by rapid pressure changes could not be reasonably
172 corrected, regardless of the sign or magnitude of the data.

173 **2.2 Supporting measurements**

174 The mixing ratios of a large number of non-methane organic compounds (NMOCs) were
175 measured by Trace Organic Gas Analyzer (TOGA) (Hornbrook et al., 2011a) and Proton-
176 transfer-reaction mass spectrometry (PTR-MS) (Karl et al., 2003; de Gouw and Warneke,
177 2007). The surface area density of fine particles was obtained by the measurement of a
178 Scanning Mobility Particle Sizer (SMPS), under the assumption of preface sphere of aerosol
179 particle. The photolysis frequencies were determined by a Charged-coupled device Actinic
180 Flux Spectroradiometer instrument (CAFS) (Shetter et al., 2002). The mixing ratios of HO_x
181 and RO_2 radicals were measured by a method based on selected-ion chemical-ionization mass
182 spectrometry (SICIMS) (Hornbrook et al., 2011b; Mauldin et al., 2010). The mixing ratios of

ozone and NO_x were measured by chemiluminescence instruments (Ridley et al., 2004). Meteorology parameters were provided by state parameter measurements on board the C-130.

3 Results and Discussion

3.1 General data description

Figure 2 shows the time series of HONO, NO_x, pNO_3 concentrations and the measurement altitude for five selected research flights in the Southeast U.S. during the NOMADSS 2013 summer field study. Research flight (RF) #4, RF #5 and RF #17 are race track flights in the background terrestrial areas designed to establish HONO distribution and explore HONO chemistry in background air masses. RF #11 is a race track flight designed to intercept plumes from local power plants and urban areas and explore HONO chemistry therein. All four flights were conducted in the daytime, roughly from 14:00 to 22:00 UTC (10:00 to 18:00 EDT). RF #18 is a race track flight conducted from 20:30 on July 12th to 03:30 on July 13th UTC (16:30 to 23:30 on July 12th, 2013 EDT), aiming to study the potential night-time HONO accumulation both in the PBL and the FT.

Table 2 summarizes the data statistics for HONO, NO_x and pNO_3 measurements in the PBL and the FT, and Figure 3 shows composite vertical distributions of HONO, NO_x and pNO_3 concentrations from the five flights in the Southeast U.S. during the NOMADSS 2013 summer field study. HONO, NO_x and pNO_3 concentrations show horizontal gradients in every race track flight and vary in different race track flights, reflecting the inhomogeneity of air masses in the region. However, there was no significant vertical gradient in HONO, NO_x and pNO_3 concentrations, which will be further discussed below. Except in a few power plant plumes and urban plumes mostly encountered in RF #11, most of the data is representative of background terrestrial air masses. The range of the mixing ratio of HONO is 3.1 – 34.4 pptv. The mean ($\pm$1SD) and median values of HONO concentration are 5.6 ($\pm$3.4) pptv and 4.2 pptv in the FT, and 11.2 ($\pm$4.3) pptv and 10.3 pptv in the PBL. HONO levels at ~ 4 pptv are typically found in the background FT, but high HONO concentrations up to 15.2 pptv are also observed in the elevated biomass burning plumes. HONO levels at ~ 11 pptv are representative of background conditions in the PBL. High HONO levels up to 34.4 pptv are observed in the power plant plumes and urban plumes in RF #11 (see section 3.4) and RF # 17. These measured HONO values are consistent with the range of 4 – 74 pptv in the troposphere over Northern Michigan (Zhang et al., 2009), but are significantly lower than 100 - 150 pptv in the morning residual layer over an industrial region of Northern Italy (Li et al.,

2014), where the levels of HONO precursors, such as NO_x and pNO_3, were much higher. The HONO concentrations were also consistent with the levels reported for the same region during the Southeast Nexus Experiment on the NOAA WP-3D aircraft (Neuman et al., 2016), that is, "indistinguishable from zero within the 15 parts per trillion by volume measurement uncertainty" in the background air and up to 150 pptv in the power plant plumes during the day. The lower HONO concentrations measured in the power plant plumes in this study than the daytime values reported by Neuman et al (2016) probably reflects greater dilution of smaller plumes encountered by the C-130 than by the WP-3D, as indicated by much lower NO_x levels observed, up to 1.8 ppbv vs up to 60 ppbv.

The range of the mixing ratio of NO_x is from several pptv to around 1.8 ppbv. The mean ($\pm$1SD) and median values of NO_x concentration are 94 ($\pm$53) pptv and 92 pptv in the FT, and 316 ($\pm$182) pptv and 279 pptv in the PBL. The mixing ratios of NO_x are mostly between 50 - 150 pptv in the background conditions in the FT and between 200 - 500 pptv in the background conditions in the PBL. Similar to HONO, high values of NO_x also occur in the urban and power plant plumes in the PBL (up to 1.8 ppbv) and in the biomass burning plumes in the FT (up to 0.6 ppbv).

Fewer measurement data points are available for pNO_3, compared to those for NO_x and HONO, due to air bubble formation in the flow cell of the pNO_3 system, especially at high altitudes. The range of the mixing ratio of pNO_3 is from 3 pptv to 186 pptv, with the mean ($\pm$1SD) and median values of 35 ($\pm$39) pptv and 15 pptv in the FT, and 76 ($\pm$45) pptv and 66 pptv in the PBL. The pNO_3 levels were highly variable in both the FT and the PBL. In the FT, the pNO_3 levels were often under 10 pptv, but high concentrations up to 179 pptv were also observed in elevated biomass burning plumes. In the PBL, high pNO_3 levels were sometimes observed in relative clean conditions; whereas, low pNO_3 levels were observed in high HONO and NO_x power plant plumes. Both the N(V) level (= [HNO_3] + [pNO_3]) and the partitioning between HNO_3 and pNO_3 seem to play roles in determining the pNO_3 level.

3.2 HONO contribution from ground-level sources

There are several ground-level HONO sources that may contribute to the HONO budget in the overlying atmosphere. They include anthropogenic sources, such as power plant and automobile emissions (Li et al., 2008b; Neuman et al., 2016), and natural processes, such as soil emission (Maljanen et al., 2013; Oswald et al., 2013; Su et al., 2011), heterogeneous

reactions of NO_2 (Acker et al., 2006; George et al., 2005; Ndour et al., 2008, 2009; Ramazan et al., 2006) and surface HNO_3 photolysis (Ye et al., 2016b; Zhou et al., 2003,2011). Since HONO photolytic lifetime is relatively short, e.g. 8 - 16 min in RF #4, RF #5, RF #11 and RF #17, a steep negative vertical gradient of HONO concentration would be expected if a significant contribution originated from the ground. The lack of a significant vertical gradient in the measured HONO concentrations (Fig. 3a) thus suggests that the ground contribution is either limited to the shallow layer of the boundary layer near the ground, below the C-130 lowest flight altitude of 300 m, or small relative to the *in situ* production of HONO in the air column (Ye et al., 2017).

To further examine the potential HONO contribution from the ground sources, vertical profiles of HONO, NO_x, and pNO_3, are compared with those of potential temperature (K) and isoprene measured, for example, in the first race-track of RF#4 from 11:00 – 12:15 LT (Fig. 4). The PBL height (H) was ~ 1200 m as indicated by the constant potential temperature (Fig. 4e). The vertical distribution of isoprene originating from the ground can be expressed with the following equation (Eq.1):

$$\ln\left(\frac{C}{C_0}\right) = -\frac{k\tau}{H}h = -\frac{h}{h^*} \qquad \text{(Eq. 1)}$$

where C_o and C are its concentrations near the ground and at the altitude h, k is the pseudo-first order degradation rate constant, τ is the average mixing time in the PBL, and h^* (= $H/(k\tau)$) is its characteristic transport height within one degradation lifetime of isoprene. According to the best fit of (Eq.1) to the observed isoprene data (Fig. 4d), its characteristic transport height h^* is estimated 692 m for isoprene. Assuming isoprene is mainly oxidized by the OH radical whose average concentration was found to be ~ 3×10^6 mole cm^{-3} in the PBL in the Southeast U.S. during the NOMADSS study (Kaser et al., 2015), the pseudo-first order degradation rate constant of ~ 3.0×10^{-4} s^{-1} (or the degradation rate of ~ 0.93 h^{-1}) is determined for isoprene. Based on a boundary layer height of ~1.2 km (Fig. 4e), an average PBL mixing time τ is estimated to be ~1.6 h between 11:00 – 12:15 LT of RF #4. With a photolytic lifetime of ~ 11 min for HONO, about 11% of the HONO originated from the ground level is expected to reach the altitude of 300 m, the lowest flight altitude of the C-130 aircraft between 11:00 – 12:15 LT in RF #4. The vertical mixing of the PBL is enhanced from the morning to the afternoon, as the ground surface is heated by solar radiation gradually during the day. The average mixing time in the PBL is reduced from ~ 3 h in the morning, to ~ 1.5 h around noontime, and to ~ 30 min in the afternoon, determined from isoprene gradients from RF #4, #5 and #17. About 50% of the ground emitted HONO could survive

and be transported to lower measurement altitudes. Again, if this ground source contribution was significant, the HONO concentration profile should exhibit a significant gradient, probably more pronounced than that of isoprene due to its shorter lifetime. The lack of such a vertical HONO gradient in the measured HONO concentration profiles (Fig. 3a) suggests that the contribution from ground HONO sources to the observed HONO concentrations in the PBL above 300 m is insignificant.

The NO_x level was $\leqslant$ 0.5 ppbv in the PBL over the Southeast U.S. (Figure 2), excluding the power plant plumes. Based on an upper limit HONO/NO_x ratio of 0.05 for urban atmosphere at ground level (Villena et al., 2011), the initial HONO concentration would be $\leqslant$ 25 pptv in the source air mass on the ground level. With a transport time of $\geqslant$ 0.5 h, i.e., $\geqslant$ 3 times of the HONO photolysis lifetime, the contribution from the ground HONO source would be $\leqslant$ 1 pptv. This analysis supports the conclusion that contribution of surface HONO source to the PBL HONO budget is insignificant.

3.3 Daytime HONO chemistry in low NO_x areas

After removing the data measured in the urban and power plant plumes, the daytime HONO concentrations are within the range of 5 - 15 pptv throughout the PBL in the background terrestrial areas in the five race-track research flights. Photolysis is the dominant sink for HONO, with a photolysis lifetime of 8 - 16 min during these four daytime flights (RF #4, RF #5, RF #11, and RF #17). Therefore, there must be a significant volume HONO source, up to 173 pptv h^{-1}, within the air mass to sustain the observed HONO concentrations.

Both NO_x and pNO_3 are potential HONO precursors in the air column. While HONO correlates moderately with NO_x (r^2 = 0.45, Fig S1), it only weakly correlates with pNO_3 (R^2 = 0.17). It appears at first that NO_x is a more important HONO precursor than pNO_3. However, the detailed analysis below suggests that NO_x is only a minor precursor to the observed HONO, and photolysis of pNO_3 is the major *in situ* HONO source.

The upper limit of photo-stationary state HONO concentration ($[HONO]_{pss}$) was calculated using Equation 2 that takes into account all the known HONO source contributions from NO_x-related reactions, including gaseous reactions of OH and NO (R-1), excited NO_2 (NO_2*) and water vapor (R4) (Carr et al., 2009; Li et al., 2008a), NO_2 and the hydroperoxyl-water complex ($HO_2 \cdot H_2O$) with an upper limit HONO yield of 3% (R5a)(Li et al., 2014; Ye et al., 2015), and heterogeneous reaction of NO_2 on aerosol surfaces (R2) using an upper limit

uptake coefficient of 10^{-4} reported in the literature (George et al., 2005; Monge et al., 2010; Ndour et al., 2008, 2009; Stemmler et al., 2006, 2007):

$$[\mathrm{HONO}]_{pss} = \frac{k_{-1}[\mathrm{NO}][\mathrm{OH}]+k_4[\mathrm{NO_2}^*][\mathrm{H_2O}]+\alpha k_5[\mathrm{NO_2}][\mathrm{HO_2{\bullet}H_2O}]+k_2 S_{aerosol}[\mathrm{NO_2}]}{J_{HONO}+k_{OH-HONO}[\mathrm{OH}]} \quad \text{(Eq. 2)}$$

where $S_{aerosol}$ is the aerosol surface area density. It should be noted that the upper limit values of rate constants were used in the calculation to avoid the underestimation of $[\mathrm{HONO}]_{pss}$ value. Under typical daytime conditions in the PBL with the median measured values of reactants, the upper limit $[\mathrm{HONO}]_{pss}$ value is less than 2 pptv, much lower than the median measured HONO concentration of ~ 11 pptv. Figure 5a shows the relationship ($r^2 = 0.40$) between the photolytic HONO loss rate with the sum of HONO production rates from all the NO_x-related reactions calculated with upper-limit reaction rate constants. A slope of about 0.19 indicates that the contribution from these NO_x-related reactions to the volume HONO source is minor in the background troposphere, despite the good correlation between HONO and NO_x.

Photolysis of HNO_3 on surfaces has been found to proceed at a much higher rate than in the gas phase (Baergen and Donaldson, 2013; Du and Zhu, 2011; Ramazan et al., 2004; Ye et al., 2016b; Zhou et al., 2003; Zhu et al., 2008), with HONO as the major product on environmental surfaces (Ye et al., 2016a, 2017). Furthermore, photolysis of particulate nitrate has been found to be the major daytime HONO source in the marine boundary layer (Ye et al., 2016b). To examine the role of particulate nitrate as a potential HONO source in the troposphere, aerosol samples were collected and used in the light-exposure experiments to determine the photolysis rate constants for particulate nitrate in the laboratory. The determined pNO_3 photolysis rate constant ($J^N_{pNO_3}$) varies over a wide range, from 8.3×10^{-5} s^{-1} to 3.1×10^{-4} s^{-1}, with a median of 2.0×10^{-4} s^{-1} and a mean (± 1 SD) of 1.9 (± 1.2) $\times 10^{-4}$ s^{-1}, when normalized to tropical noontime conditions at ground level (solar zenith angle = 0 °), and the average HONO to NO_2 relative yield is 2.0 (Ye et al., 2017). Figure 6b shows the relationship between the photolytic HONO loss rate (J_{HONO}×[HONO]) and the volume HONO production rates from pNO_3 photolysis (2/3 × J_{HNO_3} × [pNO_3]). The median $J^N_{pNO_3}$ of ~ 2.0×10^{-4} s^{-1} was used to calculate the ambient J_{pNO_3} by scaling to J_{HNO_3}:

$$J_{pNO_3} = J^N_{pNO3} \times \frac{J_{HNO_3}}{7.0\times10^{-7}\ s^{-1}} \quad \text{(Eq. 3)},$$

where J_{HNO_3} is the photolysis rate constant of gas-phase HNO_3 calculated from light intensity measurement on the C-130 aircraft, and 7.0×10^{-7} s^{-1} is the photolysis rate constant of gas-phase HNO_3 under the tropical noontime condition at ground level (solar zenith angle = 0 °). A

slope of 0.69 can be derived from Figure 5b, suggesting that pNO_3 photolysis is the major volume HONO source. However, the r^2 of 0.34 is not as strong as expected from pNO_3 photolysis being the major volume HONO source. It may be in part due to the use of a single median J^N_{pNO3}value of ~ 2.0×10^{-4} s^{-1} in the calculations of the ambient J_{pNO_3} and the production rates of HONO in Figure 5b; the actual J^N_{pNO3}values are highly variable, ranging from 8.3×10^{-5} s^{-1} to 3.1×10^{-4} s^{-1} (Ye et al., 2017). HONO source contribution from particulate nitrate photolysis in Figure 5b are thus estimates of the *in situ* HONO production rates from pNO_3 photolysis in different air masses.

HONO photolysis has been found to be an important or even a major OH primary source in the atmosphere near the ground surface (Elshorbany et al., 2010; He et al., 2006; Kleffmann et al., 2003; Villena et al., 2011; Zhou et al., 2011). However, HONO is not a significant daytime OH precursor in the background troposphere away from the ground surface. Based on the measurement results in this study, the mean (± SD) contribution of HONO photolysis to the OH source budget is 53 (± 21) pptv h^{-1} in the PBL and 44 (± 26) pptv h^{-1} in the FT (Table S1), respectively, less than 10% of the OH production contributed by O_3 photolysis. Since HONO is mainly produced from photolysis of particulate nitrate, it becomes an important intermediate product of a photochemical renoxification process recycling nitric acid and nitrate back to NO_x. The regenerating rate of NO_x of about 38 (± 23) pptv h^{-1} via pNO_3 photolysis (Table S1) is equivalent to an air column NO_x source of ~ 2.3×10^{-6} mol m^{-2} h^{-1} in the 1.5- km PBL, a considerable supplementary NO_x source in the low-NO_x background area.

It should be pointed out that particulate nitrate is in a dynamic equilibrium with gas-phase HNO_3, the later accounts for a larger (or even dominant) fraction of total nitrate (pNO_3+HNO_3) and is photochemically inert. The overall photolysis of pNO_3+HNO_3 would be much slower than indicated by J_{pNO3}. In addition, oxidation of NO_x via several reactions will replenish the pNO_3+HNO_3 reservoir. The results reported here and in earlier papers (Reed et al., 2017; Ye et al., 2016a) suggest that there is a rapid cycling in reactive nitrogen species in the low-NO_x atmosphere, sustaining the observed levels of HONO and pNO_3.

3.4 HONO chemistry in plumes

One of the objectives of RF #11 was to study the chemistry of HONO in urban and coal fired power plant plumes. The arrows and corresponding labels in Figure 2 indicate the urban plumes (U1-U3) and power plant plumes (P1-P4). CO and benzene were used to identify

influence from urban plumes, SO_2 to identify influence from power plant plumes, and acetonitrile to identify the influence of biomass burning plumes (Fig. S2). The plumes U1-U3 were identified as urban plumes from cities of Birmingham (U1, U3) and Montgomery (U2) in Alabama, respectively, and plumes P1-P4 were identified as power plant plumes from power plants in Monroe county (P1-P3) and Putman county(P4) in Georgia, respectively. The influence of biomass burning plumes was negligible as acetonitrile concentration was low and stable. The power plant plumes were generated from high-intensity point sources, and thus had features of narrow but high peaks of both HONO and NO_x concentrations in the time-series plots (Figs. 2, 6 & S2). In contrast, the urban plumes were generated from area sources and thus were shown as broad peaks of HONO and NO_x in the time-series plot with low levels of NO_x (mostly below 500 pptv) (Figs. 2, 6 & S2). There were a few sharp but small NO_x peaks within the broad urban plumes, reflecting the contributions of some point sources in the urban areas. The observed HONO/NO_x ratio was 0.019 (± 0.004) in the power plant plumes (e.g., P4) and 0.057 (± 0.0019) in urban plumes, significantly higher than the typical HONO/NO_x emission ratio of ~0.002 in the fresh power plant plumes (Neuman et al., 2016) and $\leqslant$0.01 in automobile exhaust (Kurtenbach et al., 2001; Li et al., 2008b). The elevated HONO/NO_x ratios observed in the plumes suggest that the observed HONO was mostly produced from precursors within the air mass during the transport. Based on the distances between measurement locations from the power plants or the centres of urban areas and wind speed, the transport times of these plumes were estimated to be $\geqslant$1 h, ~ 5 times longer than HONO photolysis lifetime of 8 - 16 min, again suggesting that most of the observed HONO in the plumes was produced *in situ* within the air masses.

Figure 6 shows the time-series plot of HONO budgets within the air masses sampled by the C-130 aircraft during flight RF# 11, comparing its photolysis loss rate with its production rates from pNO_3 photolysis and from all the NO_x-related reactions combined. Photolysis of particulate nitrate appears to be the major volume HONO source in all urban plumes and in most of the power plant plumes except for plume P4 observed here. NO_x was generally more important as a HONO precursor in the power plant plumes than in the urban plumes and in low-NO_x background air masses, due to higher levels of NO_x (up to 1.8 ppb in Figs. 2c and S2), OH radical and aerosol surface density. For example, all the NO_x-related reactions combined contributed up to 52% of the total volume HONO source required to sustain the observed HONO concentration in plume P4 (Fig. 6). In fresh and larger power plant plumes encountered during the RF #7 to Ohio River Valley (X. Zhou, unpublished data),

over 20 ppb NO_x was detected, and the NO_x-related reactions, mainly the NO+OH reaction, were found to account for almost all the required HONO source strength to sustain the observed HONO, in agreement with Neuman et al. (2016). The power plant plumes undergo rapid physical and photochemical evolution during the day, such as dilution and NO_x-to-HNO_3 conversion. Thus, the relative contributions from NO_x-related reactions and particulate nitrate photolysis as HONO sources change rapidly as the plumes age.

3.5 Night-time HONO chemistry

Nighttime HONO accumulation has been widely observed at the ground level (Kleffmann et al., 2003; Oswald et al., 2015; 2008; Stutz et al., 2002, 2010; VandenBoer et al., 2013, 2014, 2015), contributed by various anthropogenic and natural HONO sources on the ground. The main objective of RF #18 was to study the night-time HONO evolution in both the nocturnal residual layer and the nocturnal FT. After sunset, the surface cooling promotes the formations of a inversion layer near the ground surface and a nocturnal residual layer above; the contribution from ground HONO sources then becomes negligible to the air masses beyond the surface inversion layer. Meanwhile, no effective HONO sinks, such as photolysis, oxidation by OH and dry deposition, exist in the nocturnal residual layer. Thus the HONO accumulation, if any, is a net contribution from dark heterogeneous NO_2 reaction on aerosol surfaces (R2).

The C-130 flew in an elongated race track pattern along a north-south direction, about 140 km from Nashville, TN (Fig. 1), alternating between the PBL (1200 m) and the FT (2500 m), from late afternoon to midnight local time (Fig. 2). In the FT, HONO and NO_x concentrations were relatively stable throughout the afternoon and the night, staying around 4 ppt and 90 pptv respectively. The lack of night-time HONO accumulation is expected from the low levels of HONO precursors, mostly NO_2, and surface area of aerosol particles in the FT (Fig. 2).

The conditions in the PBL were far more variable and complicated. There were strong horizontal gradients of NO_x, pNO_3 and HONO in the PBL, with higher concentrations at the southern end and lower concentrations at the northern end of the flight track. Back-trajectory analysis using NOAA's HYSPLIT model (Stein et al., 2015) indicates that the encountered air masses in the PBL at the southern end passed over Nashville, about 140 km northeast of the sample area, with a transport time of about 6 h (Fig. S3a), while the air masses at the northern end stayed to north of Nashville (Fig. S3b). Therefore, the anthropogenic emissions from the

metropolitan area of Nashville contributed to the higher concentrations of pollutants observed at the southern end of the flight track. There were also trends of increasing concentrations of NO_x, pNO_3 and HONO with time after the sunset (Fig. 2). This was probably a result of less dispersion and dilution of anthropogenic pollutants, including NO_x, as the PBL became more stable after sunset. Furthermore, as time progressed from late afternoon into evening and night, the air masses were less photochemically aged during the transport from the source areas, due to the decreasing solar light intensity and shorter solar light exposure time.

Because of the large spatial and temporal variations in the concentrations of HONO and its precursors in the PLB (Fig. 2), it is difficult to directly evaluate the nighttime HONO accumulation from HONO measurements alone. The concentration ratio of HONO and its dominant nighttime precursor, NO_2, can be used as an indicator of nighttime HONO accumulation. As the air masses at measurement altitude of 1200 m decoupled from the ground-level processes after sunset, the HONO production ($P(\mathrm{HONO})$) from heterogeneous NO_2 reaction (R2) on aerosol surface becomes the only HONO source, and can be expressed by the following equations (Eq. 4 and Eq. 5):

$$P(\mathrm{HONO}) = \frac{1}{4} \times \left[\frac{S}{v}\right] \times \sqrt{\frac{8RT}{\pi M}} \times \gamma \times [\mathrm{NO_2}] \quad \text{(Eq.4)}$$

$$\frac{P(HONO)}{[NO_2]} = \frac{1}{4} \times \left[\frac{S}{v}\right] \times \sqrt{\frac{8RT}{\pi M}} \times \gamma \quad \text{(Eq.5)}$$

where $\left[\frac{S}{v}\right]$ is the specific aerosol surface area density, R is the gas constant, K the absolute temperature, M the molecular weight of NO_2, and γ is the dark uptake coefficient of NO_2 leading to HONO production. The NO_2-normalized HONO accumulation over time, $\Delta\frac{[\mathrm{HONO}]}{[\mathrm{NO_2}]}$, can then be calculated by equation (Eq. 6):

$$\Delta\frac{[HONO]}{[NO_2]} \sim \frac{1}{4} \times \left[\frac{S}{v}\right] \times \sqrt{\frac{8RT}{\pi M}} \times \gamma \times \Delta t \quad \text{(Eq. 6)}$$

Assuming a dark uptake coefficient γ of 1×10^{-5} of NO_2 on aerosol (George et al., 2005; Monge et al., 2010; Ndour et al., 2008; Stemmler et al., 2006, 2007) with a $\left[\frac{S}{v}\right]$ value of ~10^{-4} m^{-1}, a relative HONO accumulation rate, $\Delta\frac{[\mathrm{HONO}]}{[\mathrm{NO_2}]}/\Delta t$ of ~0.0003 h^{-1} is estimated using the equation (Eq. 6), equivalent to a HONO accumulation of 0.13 pptv hr^{-1} at a constant NO_2 concentration of 400 pptv. Such a low HONO accumulation rate is below our measurement detection limit. Indeed, the calculated HONO to the NO_x ratio using the measurement data stayed almost unchanged with time (Fig. 7), well within the observational variability after the

sunset, suggesting no significant volume production of HONO in the nocturnal boundary layer.

4 Conclusions

Substantial levels of HONO existed during the day in both the PBL (median ~ 11 pptv) and the FT (median ~ 4 pptv) over the Southeast U.S. during the NOMADSS 2013 summer field study. The ground HONO sources did not significantly contribute to the HONO budget in the PBL above the minimum measurement heights of 300 m. HONO budget analysis suggests that photolysis of particulate nitrate was the major volume HONO source (~69%), while the sum of known NO_x-related reactions a minor HONO source (~19%) in the low-NO_x background air masses. HONO was not a significant daytime OH precursor in the PBL away from the ground surface; however, HONO was an important intermediate product of photolysis of particulate nitrate in the renoxification process. Up to several tens pptv of HONO were observed in power plant plumes and urban plumes during the day, mostly produced *in situ* from precursors including NO_x and pNO_3. No significant night-time HONO accumulation was observed in the nocturnal residual layer and the free troposphere, due to low levels of NO_x and specific aerosol surface area.

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711
712

713 Table 1. Measurements from the NOMADSS 2013 summer study used in this analysis.

Parameters	Instrument	Time Resolution	Detection Limit	Accuracy	References
HONO	LPAP	200 s	1 pptv	20%	(1, 2)
pNO_3	LPAP	360 s	2 pptv	30%	(1, 2, 3)
HNO_3	LPAP	20 min	2 pptv	30%	(1, 2, 3)
NO	CI	1 s	20 pptv	10%	(4)
NO_2	CI	1 s	40 pptv	15%	(4)
O_3	CI	1 s	100 pptv	5%	(4)
OH	SICIMS	30 s	$*5\times10^4$	30%	(5, 6)
HONO	DOAS	60 s	~ 30 pptv	20%	(7)
Photolysis Frequencies	CAFS	6 s		10-15%	(8)
Surface area density	SMPS/UHSAS	65 s/1 s		20%	(9)
VOCs	PTRMS	15 s		20%	(10, 11)
VOCs/organic nitrates	TOGA	20 s		20%	(12)

714 *****in molecules cm^{-3}

715 LPAP: long-path absorption photometric (LPAP) systems

716 CI: 4-channel chemiluminescence instrument

717 SICIMS: selected-ion chemical-ionization mass spectrometer

718 DOAS: Differential Optical Absorption Spectroscopy

719 CAFS: Charged-coupled device Actinic Flux Spectroradiometer

720 SMPS: Scanning Mobility Particle Sizer

721 UHSAS: Ultra-High Sensitivity Aerosol Spectrometer

722 PTRMS: Proton Transfer Reaction Mass Spectrometry

723 TOGA: Trace Organic Gas Analyzer

724 References: (1) Zhang et al., 2012; (2) Ye et al., 2016b; (3) Huang et al., 2002; (4) Ridley et
725 al., 2004; (5) Hornbrook et al., 2011b; (6) Mauldin et al., 2010; (7) Platt and Stutz, 2008;
726 (8) Shetter et al., 2002; (9) Flagan, 2011; (10) Karl et al., 2003; (11) de Gouw and
727 Warneke, 2007; (12) Hornbrook et al., 2011a.

Table 2. Data statistics for HONO, NO_x and pNO_3 measurements both in the PBL and the FT from the five Southeast U.S. research flights during the NOMADSS 2013 summer field study. The statistics analysis is based on 1-min NO_x data, 3-min HONO data and 6-min pNO_3 data.

		HONO, pptv	**NO_x, pptv**	**pNO_3, pptv**
PBL	Range	3.1 – 34.4	81 – 1774	9 – 186
	Mean ± SD	11.2 ± 4.3	316 ± 182	76 ± 45
	Median	10.3	279	66
	N	356	904	121
FT	Range	1.3 - 15.2	<10 – 582	3 – 179
	Mean ± SD	5.6 ± 3.4	94 ± 53	35 ± 39
	Median	4.2	92	15
	N	157	655	46

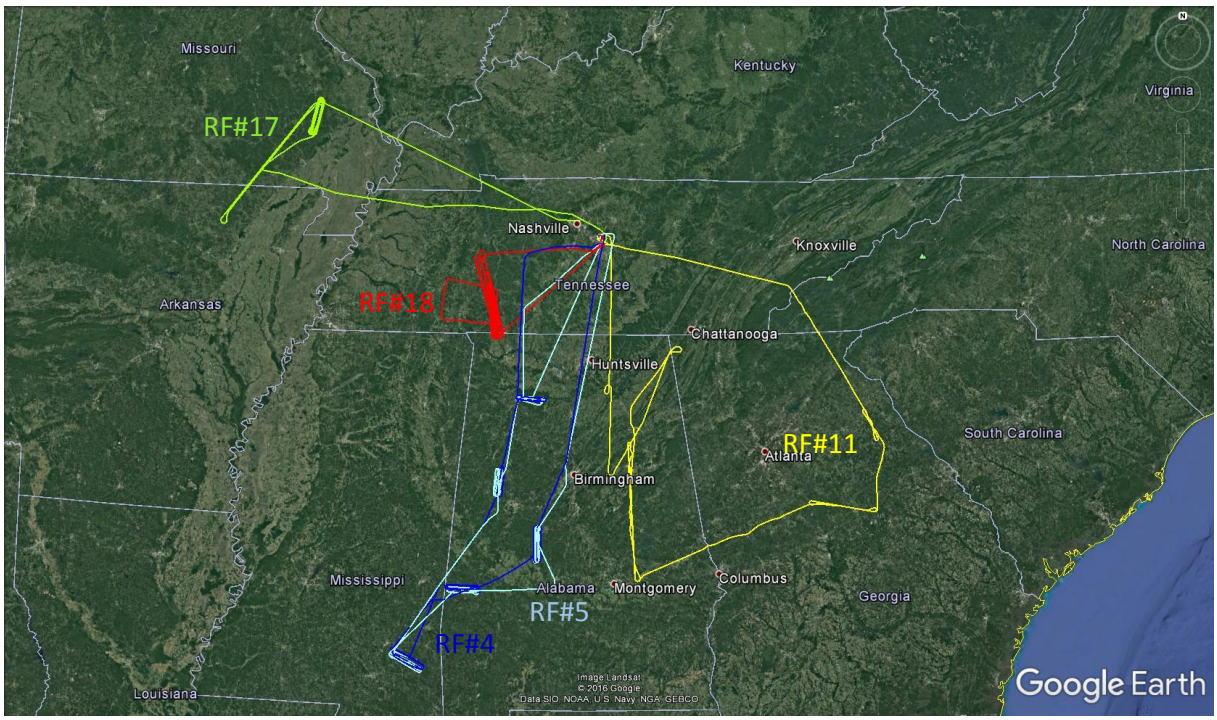
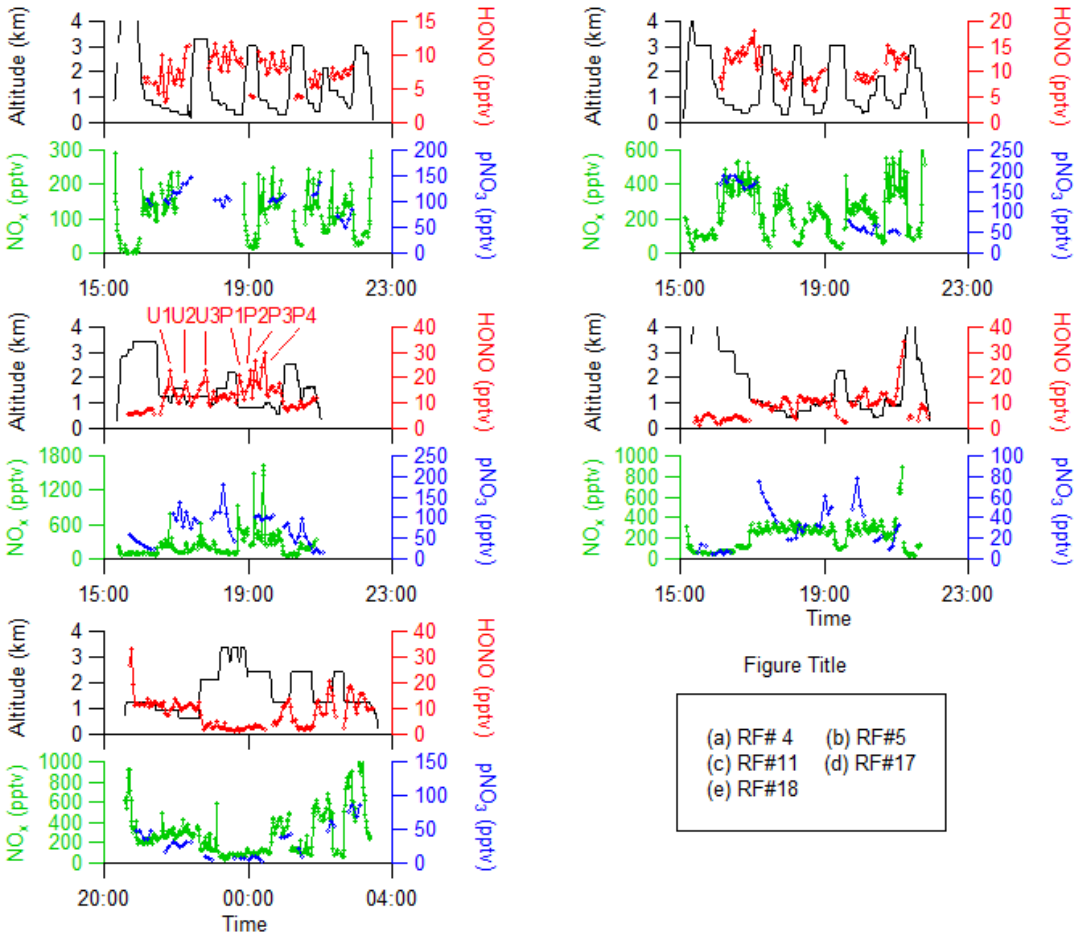


Figure 1. Flight tracks in the Southeast US during the NOMADSS 2013 summer study. The flight start time and end time in UTC (= EDT+4) are: RF#4 (blue): 15:12 and 22:30, June 12, 2013; RF#5 (light blue): 15:04 and 21:52, June 14, 2013; RF#11 (yellow): 15:20 and 21:02, June 29, 2013; RF#17 (green): 15:07 and 21:57, July 11, 2013; RF#18 (red): 20:32, July 12, 2013, and 03:37, July 13, 2013.



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741

742 Figure 2. Time series of altitude, HONO, NO_x and pNO_3 in five flights (RF #4, RF #5, RF
743 #11, RF #17 and #18) in the Southeast US during the NOMADSS 2013 summer study. In RF
744 #11, the plumes U1 and U3 were from Birmingham, AL; the plume U2 was from
745 Montgomery, AL; the plumes P1-P3 were from a power plant in Monroe County, GA; and the
746 plume P4 was from a power plant in Putnam country, GA.

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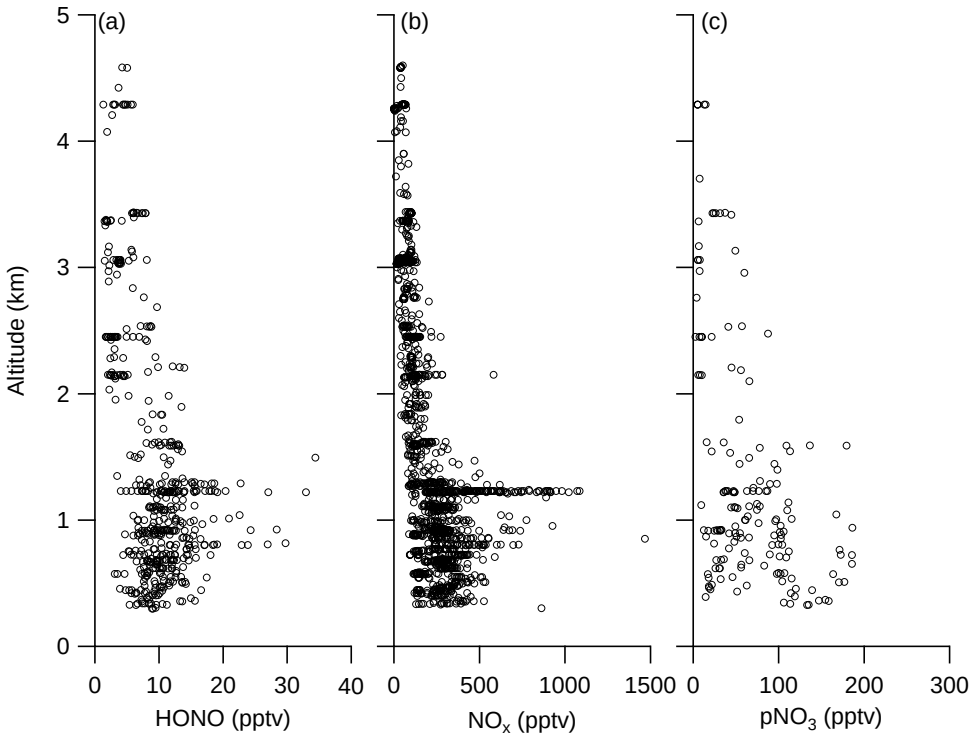

748

749 Figure 3. Vertical distributions of concentrations of HONO (a), NO_x (b), and pNO_3 (c) in the
750 five selected flights in the Southeast US during the NOMADSS 2013 summer study.

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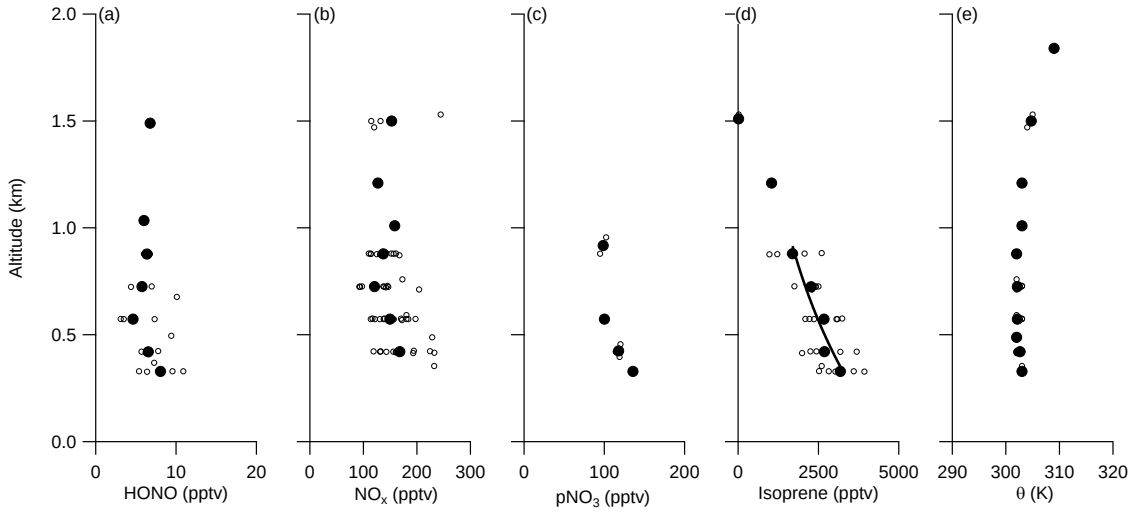
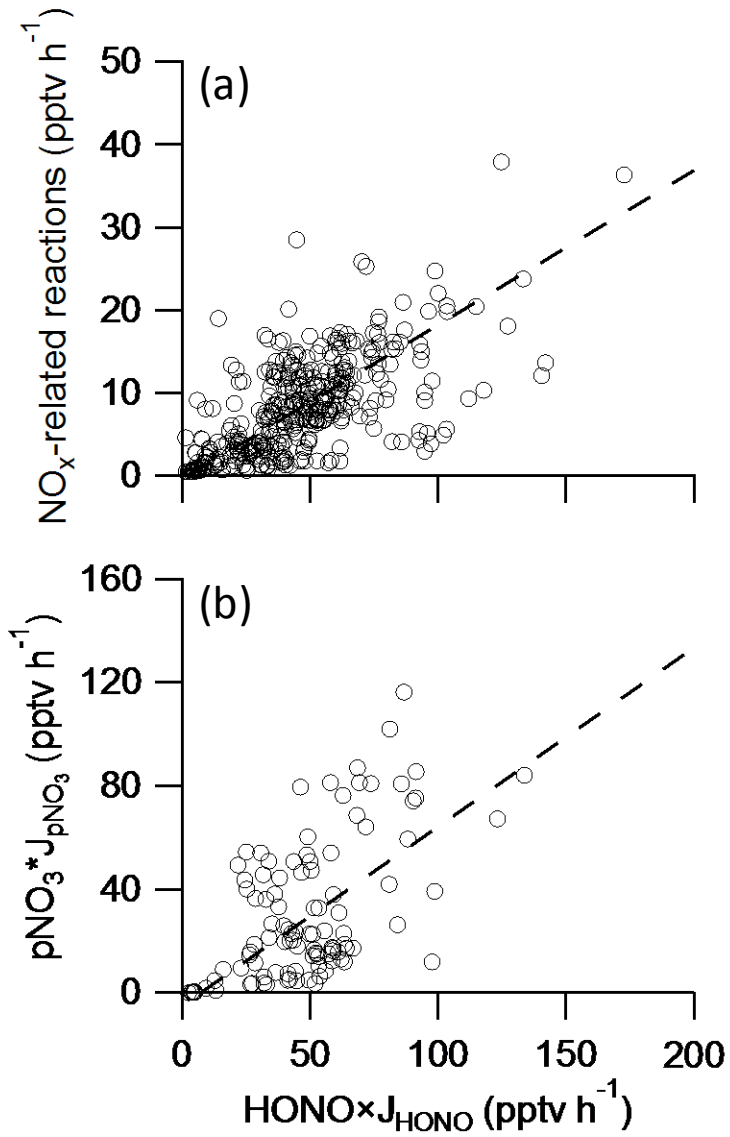


Figure 4. Vertical distributions of concentrations of HONO (a), NO_x (b), pNO_3 (c), isoprene (d) and potential temperature (e) in the PBL during the first race-track of RF#4 from 11:00 – 12:15 LT (16:00 – 17:15 UTC), June 12, 2013. The small open circles represent the 1-min data points, the large solid circles the mean values for each race-track measurement altitude. The line in (d) is the best fit of (Eq. 1) to the isoprene data: $C = 4700e^{-h/0.895}$, $r^2 = 0.93$.



760

761 Figure 5. Correlation analysis of main HONO sink ("HONO×J_{HONO}") with contribution from
762 NO_x related reactions (a) and with contribution from particulate nitrate photolysis,
763 pNO_3×J_{pNO3} (b) in the southeast US during the NOMADSS 2013 summer study. The line
764 represents the Deming least-squares regression (Wu and Yu, 2018) (r^2=0.40, intercept = -0.51
765 and slope = 0.19 for Figure 6a; r^2=0.34, intercept = -5.0 and slope = 0.69 for Figure 6b).

766

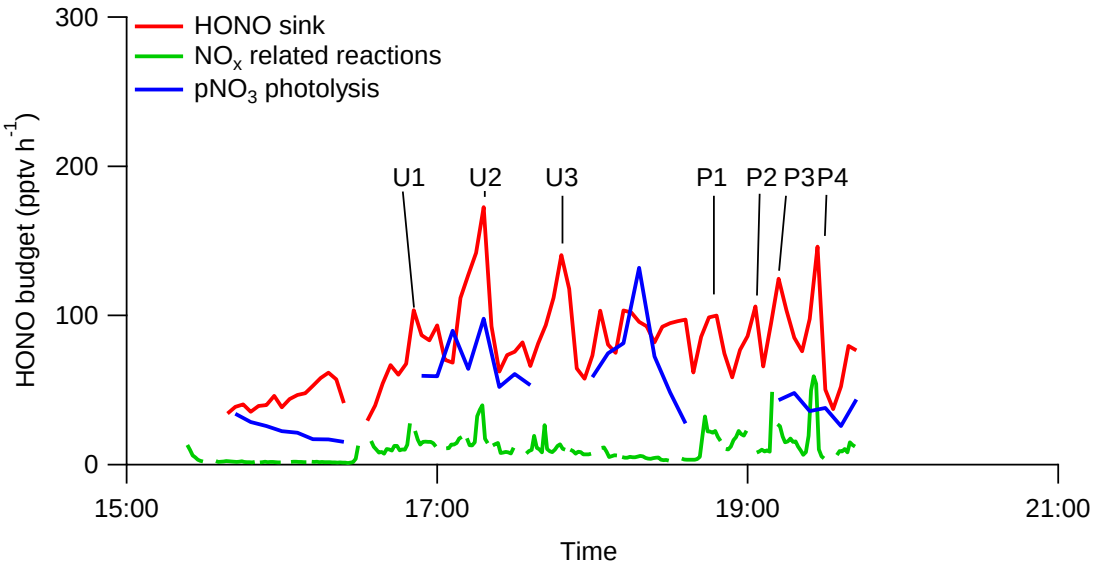
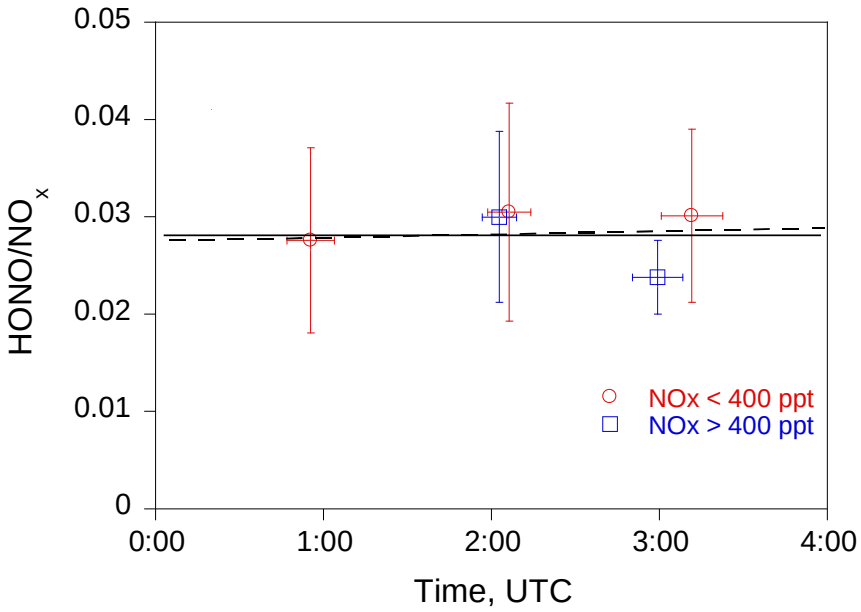


Figure 6. HONO budget analysis in RF #11 in the Southeast US during the NOMADSS 2013 summer study. "HONO sink" is the HONO loss rate contributed by photolysis and the reaction of HONO with OH radicals, "NO_x related reactions" is the sum of HONO production rates from all known NO_x reactions, and "pNO_3 photolysis" is the HONO production rate from photolysis of pNO_3. The calculations are based on 1-min NO_x data, 3-min HONO data and 6-min pNO_3 data.



776

777 Figure 7. The evolution of $HONO/NO_x$ ratio in the nocturnal boundary layer during the
778 RF#18. The red circles and blue squares are the median $HONO/NO_x$ values under the
779 conditions of $NO_x \leq 400$ pptv and $NO_x > 400$ pptv, respectively. The horizontal bars indicate
780 the averaging time periods and the vertical bars the one standard deviation of $HONO/NO_x$
781 ratios. The solid line is the least squared fit to the data, and the dashed line indicates a slope
782 of 3×10^{-4} hr^{-1}. The sunset time at the sampling location was 0:40 UTC.

783