1 **Response to comments by Forrest M. Mims III.**

2 We thank Mr. Mims for his good comments, which have helped to improve our manuscript.

3 Comment by reviewer

4 While the results in this paper are certainly intriguing, there are significant differences 5 between the instrument employed by the authors and the TOPS instrument employed by 6 Mims and Mims. The author's instrument uses filters having a FWHM bandpass of 10 7 nm, while TOPS has filters with a 5-nm bandpass FWHM. TOPS also measured column 8 ozone at 300nm and 305 nm, which is much more sensitive to ozone variations than the 9 wavelengths used by the authors. TOPS is also a direct sun instrument that can provide 10 measurements in a few seconds, while the author's instrument is a full-sky C1 device 11 with an exceptionally long 2-minute scan time. As we have learned from comparisons 12 with an EPA Brewer placed at our site, the much faster scan time provided by TOPS 13 provides higher resolution results and avoids errors caused by aerosol changes that can 14 occur during minute-duration scans. Moreover, TOPS often detects subtle changes in the 15 ozone column missed by Dobson and Brewer instruments, which both require 16 considerably more time for an ozone measurement. Before our findings are ruled out by 17 this paper, we feel that the authors should point out the very significant instrumental 18 differences, especially the filter wavelengths, the filter bandpasses and the time required 19 per scan. In each of these cases, TOPS offers superior performance when compared with 20 their instrument. Thus, the findings of subtle waves in the ozone layer by TOPS cannot 21 be so quickly discounted by this paper. I close by observing that TOPS uncovered a calibration drift in the Nimbus-7 Total Ozone Mapping Spectrometer (TOMS) (Satellite 22 23 Monitoring Error, Nature 361, 1993). TOPS evolved into Microtops and then Microtops 24 II. All these instruments provide results in close agreement with Brewers and Dobsons at 25 the Mauna Loa Observatory. Thank you for considering the points made herein. Forrest 26 M. Mims III fmimsiii@yahoo.com Interactive comment on Atmos. Chem. Phys. 27 Discuss., https://doi.org/10.5194/acp-2018-1048, 2018.

1 Authors' Response

2 As described in Section 5.2. of the manuscript, the total ozone column (TOC) was 3 derived from measurements of response-weighted global irradiance measured by the 4 GUVis-3511. Specifically, ratios of measurement at 340 and 305 nm were compared with 5 similar ratios in a look-up table that was calculated with a radiative transfer model as a 6 function of SZA and TOC. The look-up table was calculated by taking the response 7 functions of the instrument (Fig. 3 of manuscript) and observing conditions (e.g., aerosol 8 optical depth, (AOD)) into account. The method of calculating TOC from measurements 9 of global irradiance (instead of direct irradiance as it is typically done for Dobson, 10 Brewer, TOPS, and Microtops instruments) was first proposed by Stamnes et al. (1991). 11 We found that the accuracy of TOCs derived from global irradiance is similar to that of 12 data from Dobson instruments or satellite (TOMS, OMI) observations if the look-up table 13 takes local conditions into account (ozone profile, albedo, elevation, etc.) (Bernhard et 14 al., 2005b). For example, this study uncovered systematic errors in Dobson 15 measurements associated with approximations in the standard Dobson retrieval method, 16 which subsequently helped to better understand the limitation of Dobson measurements. 17 We have further validated the method for GUV instruments (Bernhard et al., 2005a). 18 Based on this work we believe that our TOC measurements are not of inferior quality 19 compared to TOPS measurements and provide further evidence below.

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21 Ability to detect small changes in TOC

As described in the manuscript (P23, L26ff.), our data do not indicate oscillations in TOC that may have been triggered by bow waves from the Moon's shadow. In contrast, Zerefos et al. (2000) reported peak-to-peak variation in TOC of about 1 %. while Zerefos et al. (2007) reported a peak-to-peak amplitude of 2.0–3.5 %. Finally Mims and Mims (1993) describe a peak-to-peak amplitude of up to 5 DU (1.7 %). We show in the following that our method is capable of detecting fluctuations of this magnitude.

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By analyzing the values of our ozone look-up table, we determined that a 1 % change in the ratio of the response-weighted global irradiance at 340 and 305 nm results in a TOC change of 1.1 DU at the start of the eclipse. At the end of the eclipse (when the SZA is

1 smaller) a 1 % change in the ratio leads to a change of 1.6 DU. As described in the 2 manuscript, ozone calculations are based on the average of 45 seconds (not 2 minutes as 3 stated by the reviewer) of global spectral irradiance measurements that are sampled at 1 4 Hz. By calculating the standard deviation of these samples and using standard error 5 propagation, we calculated an uncertainty (confidence interval of 95 %) for the ratio of 6 measurements at 340 and 305 nm of 0.13 % for the start and 0.06 % for the end of the 7 eclipse. By combining these uncertainty estimates with the sensitivity of the TOC to 8 changes in this ratio, we determined that our measurements are able to detect changes in 9 ozone of 0.14 DU at the start and 0.10 DU at the end of the eclipse. Since the average 10 TOC on 21 August 2017 was about 298 DU, these absolute changes translate to relative 11 changes of 0.05 % and 0.03 %, respectively. Our method is therefore well capable to 12 detect changes of the magnitude of 1 to 3.5 % reported by Zerefos et al. (2000, 2007) and 13 Mims and Mims (1993).

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15 We agree with the reviewer that measurements at 300 nm are more sensitive to changes 16 in ozone than measurements at 305 nm. However, measurements at 300 nm are also 17 noisier than measurements at 305 nm because the irradiance at 300 nm is smaller than 18 that at 305 nm by factors of 10 (end of eclipse) to 30 (start of eclipse). Moreover, small 19 errors in the characterization (e.g., center wavelength) of the filters functions have a 20 larger effect at shorter wavelength. Since we do not have access to a TOPS or Microtops, 21 we cannot determine whether these instruments are really superior to the GUVis-3511 in 22 determining the TOC as the reviewer asserts. (Of note, the shortest wavelength of a Dobson is 305.5 nm and if measurements at 300 nm would be of great advantage, these 23 24 instruments would likely use a shorter wavelength.) In any case, our analysis above 25 illustrates that our instrument is sensitive enough for detecting bow-waved induced TOC 26 variations of the proposed magnitude.

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28 Effect of aerosols on ozone retrieval

29 The look-up table for TOC retrievals was calculated with the same model parameters that 30 were used to convert response weighted irradiance measurements to spectral irradiance as 31 described in the Supplement to this publication. Specifically, aerosol extinction was 32 parameterized with Ångström's turbidity formula by setting the Ångström coefficients α

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1 and β to 2.10 and 0.0394, respectively. These values were derived from the instrument's measurement of direct solar irradiance at 19:00:30 after the end of the eclipse. At the start 2 3 (16:02:58; see Fig. 8b of manuscript), the Ångström coefficients were 1.96 and 0.0570, 4 respectively. To quantify the effect of changing aerosol conditions on TOC retrievals, we 5 modeled the global spectral irradiance for 16:02:58 using either the Ångström 6 coefficients used for the TOC look-up table or the coefficients applicable to this time. 7 The difference in AOD changed the ratio of global spectral irradiance at 340 and 305 nm 8 by only 0.12 %. Using the same sensitivity factor discussed above, we determined that 9 the resulting bias in TOC is 0.13 DU or 0.04 %.

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11 Limb-darkening corrected AODs (Fig. 8a), suggest that the largest AOD during the 12 eclipse occurred at 16:52 when the Ångström coefficients were 1.95 and 0.0788, 13 respectively. (This estimate is somewhat uncertain due to the uncertainty of the limb-14 darkening correction.) We calculated the aerosol effect in a similar way as before and 15 conclude that the elevated AOD at 16:52 increases the retrieved TOC values by 1.5 DU 16 or 0.5 %. Again, this value is considerably smaller than the fluctuations reported by 17 Zerefos et al. (2000, 2007). Yet, aerosols can explain 1.5 DU of the 5.0 DU increase in 18 limb-darkening corrected TOC measurements that can be seen in Fig 8a of the 19 manuscript between 16:00 and 17:00.

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We also like to point out that TOC retrievals from global irradiance is less sensitive to changes in AOD than retrievals using the direct beam (such as those utilized by the TOPS instruments) because photos that are removed from the direct beam by aerosols contribute to the global irradiance. So direct measurements are not necessarily better.

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26 <u>Sampling frequency</u>

TOC measurements discussed in the manuscript are available at a rate of one value every 28 2 minutes. It is therefore not possible to determine whether the moon's shadows resulted 29 in fluctuations in TOC on a shorter time scale. However, oscillations reported by Zerefos 30 et al. (2000) had a period of 20 minutes and those reported by Zerefos et al. (2007) had 31 periods ranging between 30 and 40 minutes. Zhang et al. (2017) reported ionospheric 32 bow waves, which manifested themselves as electron content disturbances, with a period

1 of about 25 minutes during the solar eclipse of 21 August 2017. If these bow waves had 2 also affected the ozone layer, these oscillations should have been detectable with our 3 sampling frequency. Finally, Mims and Mims (1993) report that TOC measurements 4 taken during the total solar eclipse of 11 July 1991 show three fluctuations, which began 5 700 seconds after the third contact with durations of 378, 270 and 432 seconds, 6 respectively. It is curious that these oscillations had a much shorter duration than those 7 reported by Zerefos et al., (2000, 2007) and Zhang et al. (2017). In any case, oscillations 8 of this duration should still shown up in our 2-minute data considering the low noise in our data discussed above and illustrated in Fig. 9, in particular between the 3rd and 4th 9 10 contact.

11

12 <u>Concluding remarks</u>

13 We did not "rule out" the findings of the paper by Mims and Mims (1993) as stated by 14 the reviewer. Instead, we simply stated that our data do not support the observation by 15 Zerefos (2000; 2007) and Mims and Mims (1993), and concluded that the question of 16 whether or not bow waves from the Moon's shadow can lead to variations in TOC is still 17 up for debate. We suggested that this debate could be settled by performing limb-18 darkening-corrected measurements of TOC with different instrument types during one of 19 the upcoming solar eclipses. If such measurements were to show fluctuations in TOC 20 with the same magnitude and timing, the effect of bow waves on TOC could be 21 convincingly demonstrated. We still believe that this is a reasonable path forward and 22 might stimulate future research.

23

24 **Proposed changes to manuscript**

25 The following text will be added to Sect. 5.2 of the manuscript:

"By analyzing the sensitivity of TOC to changes in the irradiance ratio and by
quantifying the noise in the measurements of the 305 and 340 nm channel we
determined that our measurements are able to detect changes in ozone of 0.14 DU
(0.05 %) at the start and 0.10 DU (0.03 %) at the end of the eclipse at 95 %
confidence level for constant aerosol conditions. The additional uncertainty in
TOC values due to changing AODs was determined to be 1.5 DU (0.5 %)."

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2	In Sect. 7.1., the following sentence:
3	"Results corrected for the LD effect (thick lines in Fig. 8a) indicate that the AOD
4	was monotonically decreasing over the period of the eclipse (from 0.41 to 0.32 at
5	319 nm and from 0.05 to 0.04 at 1018 nm) without a spurious spike near the time
6	of totality."
7	will be replaced with:
8	"Results corrected for the LD effect (thick lines in Fig. 8a) indicate that the AOD
9	was increasing between the 1^{st} contact and 16:52 (from 0.41 to 0.58 at 319 nm
10	and from 0.05 to 0.07 at 1018 nm) and then monotonically decreasing to 0.34 at
11	319 nm and 0.04 at 1018 nm at the end of the eclipse. Corrected results do not
12	have a spurious spike near the time of totality."
13	
14	In Sect. 7.2, the following sentence:
15	"(We have no explanation for the increase in TOC of about 5 DU between 16:00
16	and 17:00 other than natural variability)"
17	will be replaced with:
18	"The increase in TOC of about 5 DU between 16:00 and 17:00 can partly be
19	explained by increasing AODs over this period, which cause a high-bias in the
20	ozone retrievals."
21	
22	In Sect. 8.3, the following sentence:
23	"Bow waves with wavelengths between of 300 and 400 km and a period of about
24	25 minutes have indeed been observed during the solar eclipse of 21 August 2017
25	(Zhang et al., 2017)."
26	will be replaced with:
27	"Ionospheric bow waves with wavelengths between of 300 and 400 km and a
28	period of about 25 minutes, which manifested themselves as electron content
29	disturbances, have indeed been observed during the solar eclipse of 21 August
30	2017 (Zhang et al., 2017). If these bow waves had affected the ozone layer, these

1	oscillations should have been detectable with our sampling frequency of one TOC
2	value every two minutes."
3	
4	Also in Sect. 8.3, the following sentence:
5	"These small variations are well within the natural variability of the TOC."
6	will be replaced with:
7	"These small variations are well within the natural variability of the TOC and the
8	uncertainty of our TOC retrieval of 1.5 DU (0.5 %), which is mainly caused by
9	the effect of changing aerosols on ozone calculations."
10	
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