



1 **Seasonal and diurnal variability in air pollutants and short-lived climate forcers**
2 **measured at the Rwanda Climate Observatory**

3

4 H. Langley DeWitt¹, Jimmy Gasore^{1,3,4}, Maheswar Rupakheti², Katherine E. Potter¹,

5 Ronald G. Prinn¹, Jean de Dieu Ndikubwimana³, Julius Nkusi³, and Bonfils Safari⁴

6

7

8 ¹ Massachusetts Institute of Technology, Center for Global Change Science, Cambridge,

9 MA, USA

10 ²Institute for Advanced Sustainability Studies (IASS), Potsdam, Germany

11 ³Ministry of Education, Climate Secretariat, Kigali, Rwanda

12 ⁴University of Rwanda, Physics Department, Kigali, Rwanda

13



14 **Abstract**

15 Air pollution is still largely unstudied in sub-Saharan Africa, resulting in a gap
16 in scientific understanding of emissions, atmospheric processes, and impacts of air
17 pollutants in this region. The Rwanda Climate Observatory, a joint partnership
18 between MIT and the government of Rwanda, has been measuring ambient
19 concentrations of key long-lived greenhouse gases and short-lived climate-forcing
20 pollutants (CO_2 , CO, CH_4 , BC, O_3) with state-of-the-art instruments on the summit of
21 Mt. Mugogo (1.586° S, 29.566° E, 2590 m above sea level) since May 2015. Rwanda is
22 a small, mountainous, and densely populated country in equatorial East Africa,
23 currently undergoing rapid development but still at less than 20% urbanization. The
24 position and meteorology of Rwanda is such that the emissions transported from both
25 the northern and southern African biomass burning seasons affect BC, CO, and O_3
26 concentrations in Rwanda. Black carbon concentrations during Rwanda's two dry
27 seasons, which coincide with the two biomass burning seasons, are higher at Mt.
28 Mugogo than in major European cities. Higher BC baseline concentrations at Mugogo
29 are loosely correlated with fire radiative power data for the region acquired with
30 MODIS satellite instrument. Spectral aerosol absorption measured with a dual-spot
31 Aethalometer also varies in different seasons, likely due to change in types of fuel
32 burned and direction of pollution transport to the site. Ozone concentration was
33 found to be higher in air masses from southern Africa than from northern Africa
34 during their respective biomass burning seasons. These higher ozone concentration
35 in air masses from the south could be indicative of more anthropogenic emissions
36 mixed with the biomass burning emissions from southern Africa as Rwanda is



37 downwind of major East African capital cities in this season. During the rainy season,
38 local emitting activities (e.g., cooking, transportation, trash burning) remain steady,
39 regional biomass burning is low, and transport distances are shorter as rainout of
40 pollution occurs regularly. Thus local pollution at Mugogo can be estimated during
41 this time period. Understanding and quantification of the percent contributions of
42 regional and local emissions is essential to guide policy in the region. Our
43 measurements indicate that air pollution is a current and growing problem in
44 equatorial East Africa that deserves immediate attention.

45

46 1. Introduction

47 According to recent data collected and published by the World Bank,
48 particulate air pollution in most African countries is above the annual average
49 guideline values recommended by the World Health Organization (WHO). Despite
50 this, little scientific research has been performed on air quality in Africa. The WHO
51 reports that one in eight premature deaths globally can be linked currently to poor air
52 quality, and these deaths are concentrated in developing countries (WHO, 2013).
53 Black carbon (BC) is one of the major air pollutants emitted from Africa, mainly from
54 biomass burning as it is widespread on the continent during certain seasons. In
55 addition to affecting health, BC contributes to atmospheric heating and thus to climate
56 change (Ramanathan and Carmichael, 2008). Widespread crop fires in northern and
57 southern Africa, prevalent in boreal winter (December-January-February, DJF) and
58 austral winter (June-July-August, JJA), respectively, are known to increase aerosol and
59 ozone concentrations in this region and transported molecular and aerosol fire



60 tracers associated with elevated ozone have been measured as far as the Pacific and
61 Indian Oceans (Field et al., 2016; Real et al., 2010).

62 In addition to biomass burning, recent work has pointed to diffuse and
63 inefficient combustion emissions (DICE) from anthropogenic sources as having a
64 significant effect on air quality in sub-Saharan Africa, and densely populated areas
65 have higher DICE concentrations (Marais and Wiedinmyer, 2016). Examples of DICE
66 include agricultural burning, trash burning, cook stoves, kerosene lanterns, brick
67 kilns, charcoal making, vehicles, and diesel generators (Marais and Wiedinmyer,
68 2016). The rapid development of Africa, if not combined with development and
69 enforcement of emission regulations and concurrent air quality monitoring, could
70 lead to increased emissions of air pollutants and greenhouse gases from Africa to the
71 global atmosphere in the near future as vehicle use, industry, generator use, and
72 fossil-fuel power plants increase (Lioussé et al., 2014). Past studies have suggested
73 that pollution emitted near the equator has a larger impact on tropospheric ozone
74 than pollution emitted in other regions due to consistent sunlight and heat available
75 in the region to produce ozone (Zhang et al., 2016). Much of Africa's population is
76 near the equator, and population density is one driver of DICE emissions. Collecting
77 data on these diffuse emission sources is difficult, and continuous and highly time-
78 resolved (one hour or shorter time resolution) on-ground data is scarce in both major
79 cities and rural areas in East and Central Africa. While satellite data is beginning to
80 fill the gap of column-integrated air quality approximations, these approximations
81 often don't provide ground-level information on concentrations of pollutants, and
82 even they do, they are associated with later uncertainties. Currently, South Africa has



83 the most extensive network of air quality monitoring networks and highest number of
84 completed field campaigns of the sub-Saharan African countries (Sauvage et al., 2005;
85 Tiitta et al., 2014). The studies in South Africa have found that satellite data, as
86 currently understood, is not an ideal proxy for on-ground measurements as there can
87 be poor agreement between the two types of measurements (remote column versus
88 in situ boundary layer) (Hersey et al., 2015).

89 Due to logistics of long-term measurement and lack of in-country scientific
90 capacity, most studies completed in East and Central Africa have been short-term
91 studies focused on either near-roadway air pollution or exposure to PM and CO
92 emissions from cook stoves (Galbraith, 2014; Gatari and Boman, 2003; Kinney et al.,
93 2011; Knippertz et al., 2015; Mkoma et al., 2009; Ngo et al., 2015) These studies find
94 highly variable levels of air pollution, with BC comprising a significant percentage of
95 aerosol particles (Gatari and Boman, 2003; Kinney et al., 2011; Mkoma et al., 2009;
96 Ngo et al., 2015). The daily averaged ambient PM₁₀ (particulate matter 10
97 micrometers or less in diameter) concentrations varied in these studies from 7 to
98 over 80 $\mu\text{g m}^{-3}$. This huge spread in aerosol concentrations measured in East Africa,
99 varying due to season and location, demonstrates the need for long-term high-
100 frequency measurements to fully understand the complexity of air quality issues in
101 this region. Short-term studies do not capture seasonal changes in emissions and
102 transport, very likely significant as large-scale biomass burning emissions mix with
103 increasing urban and industrial emissions. Emission sources and source
104 characteristics are certain to be different than those in Europe or North America,



105 making in-situ data essential to increase our understanding of emissions from Africa
106 and also improve the accuracy of atmospheric models for Africa.

107 Located in east-central Africa, Rwanda is the second most densely populated
108 African country, following Mauritius. Approximately 11.6 million people lived within
109 its 26,338 square kilometers as of 2016. Urbanization remains low at 18% (World
110 Bank, 2011) and the majority of the population is involved in subsistence agriculture,
111 but this is changing quickly. The country is undergoing rapid development, averaging
112 an 8% per annum growth during 2001-2015. Rwanda's industrial sector still remains
113 undeveloped at this time and a low urbanization rate and high vehicular cost means
114 that large and continuous traffic jams are not yet an issue. However, combustion
115 emissions from other small but numerous diffuse sources, including the use of
116 biomass as the primary cooking fuel by the majority of the population, are likely to be
117 significant, particularly due to the high population density of Rwanda (Hersey et al.,
118 2015; Marais and Wiedinmyer, 2016). These incomplete combustion emission
119 sources, which emit CO and BC as well as other air pollutants, remain present and
120 steady throughout the year (Bond et al., 2013). Many of these emissions lead to
121 adverse health impacts. Lower respiratory infections are currently the largest single
122 cause of death in Rwanda, leading to 11% of all deaths in 2015 and 17% of all deaths
123 in children under five. Although these infections can't be directly traced to air quality,
124 air quality is known to affect the respiratory health of children (IHME, 2016). This
125 underscores air quality as a potential but unmeasured health risk in this region, as air
126 quality is tied to lower respiratory disease, particularly in children (WHO, 2013).



127 To mitigate environmental problems often associated with rapid growth, the
128 government of Rwanda has enacted a number of environmental regulations, including
129 a ban on plastic bags, increased taxation on the importation of older vehicles, taxation
130 of charcoal fuel within major cities, and age limits on motorcycles and buses brought
131 in to the country for commercial use. Data on the effects of these current and future
132 policies is needed for their justification and enforcement. The economic burden of
133 regulation in this low-income country needs to be weighed against the lowering of
134 health costs resulting from air quality improvement. Additionally, air quality issues in
135 this region are not just restricted to individual countries, and overall increases in
136 development and population density may be leading to a general increase in regional
137 air pollution in east-central Africa.

138 Rwanda is located in the middle of the two major seasonal biomass burning
139 regions of sub-Saharan Africa. Wide-scale biomass burning occurs to the north of
140 Rwanda during December-January-February (DJF) and to the south during June-July-
141 August (JJA). Rwanda's climate may exacerbate fire haze pollution effects, as Rwanda
142 experiences two dry seasons that occur at the same time as these two continental
143 burning seasons, making long range transport with low rainout efficiency likely.
144 Rwanda's prevalent wind direction also changes from northerly (DJF) to southerly
145 (JJA) at the same time as the large-scale biomass burning area shifts from north-
146 central Africa to southern Africa. Increase in incidence and amount of biomass
147 burning is thought to be one consequence of climate change in this region (Niang et
148 al., 2014). Southern Africa's biomass burning is also influenced significantly by human
149 activity, not just the climate (Archibald et al., 2010). Rwanda is positioned to



150 experience both large-scale (transported) haze due to fires and human activities and
151 local, diffuse emissions.

152 In addition to air quality issues, climate change (related to air pollution) may
153 also adversely affect Rwanda. The main products exported (coffee and tea), the
154 livelihood of the majority of Rwandans (agriculture), and power (currently almost
155 half of Rwanda's power is hydroelectric) are all potentially affected by climate change.
156 These issues are similar across the region. Central Africa is expected to receive
157 increased severe rainstorms, which may lead to erosion and an uptick in vector-borne
158 diseases (Niang et al., 2014). However, there is limited on-ground data on air quality
159 and climate change in Africa.

160 Therefore, in order to advance our scientific understanding of air pollution,
161 climate change, and their impacts in Africa through generation of on-the-ground data,
162 MIT and the government of Rwanda have established the Rwanda Climate
163 Observatory (RCO) to measure long-lived greenhouse gases and short-lived climate
164 forcers/pollutants in East Africa. Since May 2015, CH₄, CO, CO₂, O₃, and BC
165 concentrations have been continuously measured, and N₂O measurements were
166 added in February 2017.

167 The RCO is a part of the Advanced Global Atmospheric Gases Experiment
168 (AGAGE) network, a global network of high-frequency trace greenhouse gas
169 measurements (Prinn et al., 2000), and is the first station of its kind in Africa. Here we
170 present first results on diurnal and seasonal variations in short-lived climate
171 forcers/pollutants related to air quality, focusing on O₃, CO, and BC observed at the
172 RCO. This dataset is unique and unprecedented to the region. We attempt to explain



173 seasonal differences in observed concentrations of these pollutants and their
174 association with regional and local sources, meteorological conditions and
175 repercussions for Rwanda. This information on overall concentrations, sources, and
176 time-dependent concentration variations of these air pollutants is essential in this
177 rapidly changing area of the world to not only advance our understanding of air
178 pollution and climate change in the region but also inform future policies on air
179 pollution with sound science.

180

181 ***2. Experimental Methods: Rwanda Climate Observatory***

182 *2.1 Rwanda Climate Observatory Environment*

183 The RCO is located in the Northern Province of Rwanda, near Byangabo on the
184 summit of Mt. Mugogo (1.586° S, 29.566° E, 2590 m above sea level). Mt. Mugogo is
185 about 70 km (aerial distance) to the north-west from Kigali (population of
186 approximately 1 million), 20 km (south-west) from the next major city, Musanze
187 (population of around 100,000), and 60 km north-east from the Lake Kivu region
188 (Gisenyi, Rwanda and Goma, DRC, combined population of approximately 1 million). A
189 dirt road reaches the base of the mountain, about 500 m below the summit where the
190 RCO is located, and a diesel generator is installed on the road at the base. Inlets were
191 installed on both the roof of the Observatory (10 m above ground level) for O₃ and BC)
192 and on a Rwanda Broadcasting Authority Tower (35 m above ground level) for CO,
193 CO₂ and CH₄. There is a small Rwandan army camp adjacent to the measurement site
194 and a eucalyptus forest and a mix of agricultural fields and scattered rural houses
195 surround the immediate vicinity of the RCO (Figure 1). The high altitude and remote



196 positioning of Mt. Mugogo allows sampling of regional air masses from throughout
197 East Africa depending on prevailing meteorological conditions, as well as local
198 pollution (as the dense population but low urbanization of Rwanda means that direct
199 human influence is everywhere except within the national parks).

200

201 *2.2 Instrumentation*

202 Details on the instruments sampling at the RCO are compiled in Table 1. PM_{2.5}
203 BC (particulate matter 2.5 micrometers in diameter or less) was measured using a
204 Magee Scientific 7-wavelength Aethalometer with dual-spot technology that is able to
205 correct for filter loading artifacts (Drinovec et al., 2015). A cyclone PM_{2.5} impactor
206 was installed on the inlet to remove larger particles. Air was passed through a filter
207 once per day to collect blank data. Flow was calibrated once per year and after major
208 instrument movement and changes, while the optical performance was calibrated
209 with a neutral density filter kit once per year. Data was recorded every minute at a 5
210 liter per minute (LPM) flow rate and particles were captured on a quartz fiber filter
211 tape. The air stream was not dried and the relative humidity (RH) was not
212 controlled, which could lead to increased uncertainty during periods of high relative
213 humidity. RH recorded at the station varied by approximately 5% over the day and
214 from 60-85% monthly, depending on the season. The 880 nm channel was used to
215 calculate the concentration of BC.

216 CO mixing ratios were measured in real-time using a cavity ring-down
217 spectrometer (G2401, Picarro, USA). Sampled, laboratory, and calibration air were
218 dried with a Nafion drier inside an Earth Networks calibration box to increase the



219 accuracy of the Picarro water vapor correction (Welp et al., 2013). Three NOAA-
220 standard calibration tanks were used for calibration spanning normal ambient
221 concentrations and calibrations were performed once per day initially to check for
222 linearity of instrument's response (Gasore, 2018). An O₃ monitor (T400, Teledyne
223 Advanced Pollution Instrument, USA) was used to measure O₃. Regular checks were
224 performed using internal span and zero O₃ calibrations. Flow was calibrated two to
225 three times per year.

226 Meteorological data (ambient temperature, relative humidity, pressure, wind
227 speed, wind direction and rainfall) were collected with an automatic weather station
228 (WXT520, Vaisala, Finland). The weather station was attached to a fixed, hinged arm
229 35 m above ground level and connected to the communications tower, level with the
230 CO/CO₂/CH₄ inlet, with a 2 m clearance from the tower.

231 **3. Results and Discussion**

232 Figure 2 shows the time series of 15 minute and daily averaged BC, O₃ and CO
233 concentrations monitored at the RCO from May 2015 through January 2017. The 15
234 minute averaged data were used to further investigate daily, diurnal, weekly, monthly
235 and season variations. The diurnal, weekly, and monthly variations in concentrations
236 of each species, normalized to their average, are shown in Figure 3.

237 **3.1 Seasonal Variation in BC, CO, and O₃**

238 It has been known for some time that wide-scale biomass burning in sub-
239 Saharan Africa has a large seasonal effect on the atmosphere (Archibald et al., 2010;
240 Crutzen and Andreae, 1990). Understanding and separating these seasonal effects
241 from anthropogenic emissions can be difficult without continuous data sets both



242 during and outside of this period, especially as both biomass burning and
243 anthropogenic emissions in this region of the world emit BC, CO, and PM, and
244 anthropogenic emissions contain O₃ precursors that increase O₃ formation. Needed
245 on-ground measurements of various aerosol and gas-phase species associated with
246 fire and anthropogenic emissions in Africa are scarce, but recently the US Embassies
247 in Addis Ababa, Ethiopia, and Kampala, Uganda have begun continuously measuring
248 PM_{2.5} concentrations. The raw data is collected and reported online on the OpenAQ
249 platform (OpenAQ.org). This dataset on PM_{2.5} concentrations in major cities over
250 different seasons in this region has been valuable in gaining basic insights into the
251 seasonal characteristics of PM_{2.5} concentrations in the region (Figure 4). The PM_{2.5}
252 concentrations in both these cities showed clear seasonal patterns, though the
253 seasonal patterns differed at the two sites. Addis Ababa (Ethiopia) is much further
254 north than Rwanda and Ethiopia is in general higher in elevation than Rwanda
255 (though at 2355 m, not higher than the RCO) and closer to the Indian Ocean. In Addis
256 Ababa, the dry season is also in DJF, but measured PM_{2.5} concentrations were low
257 during this season. HYSPLIT back trajectory calculations confirmed that air masses
258 during this time of the year originated over the ocean, not from the continent. JJA is
259 the rainy season in Ethiopia; however, back trajectories confirmed that air masses
260 originating from fires over Madagascar and southern Africa were likely transported to
261 Addis Ababa at times and an enhancement of almost 20 µg m⁻³ for the JJA monthly
262 averages of PM_{2.5} was observed. Kampala, Uganda is close to Rwanda, near the
263 equator, and has a long dry season during JJA and a short dry season during DJF.
264 Rainy and dry season extrema are shown in the available Kampala PM_{2.5} data, with



265 an enhancement during February and JJA of around 15 to 25-30 $\mu\text{g m}^{-3}$, respectively,
266 above PM_{2.5} concentrations during other months. However, in Kampala during all
267 months measured, including the rainy season where little regional biomass burning
268 influence is likely, monthly averages remained above the WHO recommendations for
269 air pollution levels at daily averages of 25 $\mu\text{g m}^{-3}$ or less and, despite having a lower
270 population than Addis, were consistently higher in PM_{2.5} concentrations. South Africa
271 has the most air quality monitoring stations of any sub-Saharan African country and
272 results from these stations show a PM_{2.5} peak only in the southern burning season
273 (JJA), not surprisingly missing transported pollution from the northern (DJF) burning
274 season (Hersey et al., 2015).

275 From these data, though there are only two data points, it appears that African
276 countries near the equator may be positioned to experience six months per year of
277 transported regional fire haze, from both the northern and southern biomass burning
278 seasons. This is a concerning public health issue as equatorial Africa is densely
279 populated, which means that many people will be affected by transported pollution,
280 and the higher population density will increase the local diffuse pollution emissions
281 (e.g., cooking fires, diesel engines), exacerbating the problem of transported fire haze
282 pollution with additional locally emitted pollution.

283 BC and CO data from the RCO were examined to probe seasonal changes in air
284 pollution in Rwanda. Like Uganda, Rwanda has two rainy seasons roughly occurring
285 in March-April-May (MAM) and September-October-November (SON), and two dry
286 seasons during December- January-February (DJF) and June-July-August (JJA). This
287 generalized definition and durations of the seasons are used the purpose of



288 comparing data for multiple years. They therefore don't represent exact onset and
289 withdrawal dates of each rainy and dry season, as these dates are different each year.
290 High variations in BC concentrations can be seen in the BC time series (Figure 2 and
291 Figure 3), ranging from below 100 to above 20,000 ng m⁻³, with an average value of
292 1,700 ng m⁻³ (standard deviation: 1,600 ng m⁻³). Peak concentrations corresponded to
293 dry seasons. Spikes in BC concentrations that lasted for less than 15 minute with
294 values higher than 25,000 ng m⁻³ were removed before conducting any further
295 analysis to eliminate BC sources in the direct vicinity of the RCO. CO and O₃ mixing
296 ratios also increased during the dry seasons compared to the rainy seasons.

297 To explore the sources of BC, CO, and O₃ at the RCO, seven-day HYSPLIT back
298 trajectories were run every 6 hours using NCEP/NCAR reanalysis meteorological data
299 (Kalnay et al., 1996). This analysis provided insights on the approximate origin and
300 trajectories of air masses before arriving at RCO measured at the RCO. These
301 HYSPLIT back trajectories were separated into DJF, MAM, JJA, and SON and are shown
302 with MODIS satellite fire count data colored by fire radiative power (FRP, W m⁻²)
303 (Figure 5). The MODIS fire count data and radiative power are used strictly for
304 qualitative, not quantitative, purposes in this work. Here we observe that, as major
305 biomass burning sites moved to the north and west in DJF, transport direction was
306 also primarily northerly, and as biomass burning move to Southern Africa in JJA, the
307 prevailing wind directions were also southerly. Although Rwanda itself had few large-
308 scale fires, its geographical position and meteorology meant that it experienced
309 transported fire haze from both major burn seasons. Black carbon measured at the
310 station tracked fairly well with summed daily FRP for sub Saharan Africa (Figure 5),



311 with R values varying from 0.75-0.10 depending on the month. Daily averages of BC
312 at the station often exceeded $5 \mu\text{g m}^{-3}$, and the yearly average BC measured at the
313 station was greater than many rural measurement locations around the globe and on-
314 par with urban measurements in North America and Europe, though much lower than
315 measurements made in cities in China (Figure 6).

316 To further examine pollution transport to the RCO, the HYSPLIT back
317 trajectory geographical areas were gridded (using the R Openair package, (Carslaw
318 and Ropkins, 2012)) and merged, using date and time, with measured BC
319 concentrations and mixing ratios of O_3 and CO to generate concentration-weighted
320 back trajectories (cwt) for each season (more details on cwt available in (Hsu et al.,
321 2003; Seibert et al., 1994))(Figure 7). Trajectory time in each grid and arrival time of
322 each air mass were taken into account in this model to predict the likely source
323 regions and emission concentrations of pollutants measured at the RCO. This was
324 done to determine likely source regions of air pollution at the RCO by comparing
325 arrival times of air masses to the RCO and the time series of pollutants. This method
326 has proven fairly effective at identifying emission sources when comparing predicted
327 emission regions to emissions inventories (Lupu and Maenhaut, 2002) and is good as
328 a rough estimate of emission regions with no apriori information (Kabashnikov et al.,
329 2011). This method has low computational cost and is simple to set up, both of which
330 are important for areas with limited bandwidth or computational capacity and this
331 method can be repeated easily by in-country scientists.

332 BC and CO appeared to originate from similar areas, as expected due to their
333 overlapping sources of inefficient combustion and biomass burning. During JJA,



334 significant BC and CO appeared to originate from southern Africa and Madagascar, as
335 well as from local sources near the RCO. During DJF, the source of these pollutants
336 appeared to be much closer to the RCO, as major fires in the DRC and Uganda were
337 also closer to the station. Throughout the measurement period, but particularly DJF,
338 the Lake Kivu region also appeared to be a source of BC and CO. The Lake Kivu region
339 is densely populated and use of both cook stoves and diesel generators is common.

340 In addition to direct emissions of BC and CO, other emissions such as volatile
341 organic compounds and oxides of nitrogen from biomass burning are known to affect
342 tropospheric O₃ concentrations, as they are precursors to O₃ formation (Jaffe and
343 Wigder, 2012; Sauvage et al., 2005). It appears that such emissions likely played a role
344 in the observed seasonal increase in O₃ mixing ratios of approximately 20 ppb in DJF
345 and 25 ppb in JJA above rainy season levels at the RCO. The O₃ mixing ratio was
346 highest during the JJA dry season, while BC had approximately the same monthly
347 average concentrations in both dry seasons. This increase of about 5 ppb O₃ during
348 JJA versus DJF was potentially due to the mixing of biomass burning emissions with
349 anthropogenic emissions from east African cities such as Nairobi, Dar Es Salam, and
350 Kampala during the JJA dry season. Direct source apportionment of O₃ is difficult as it
351 formed downwind of emissions, but a mix of biomass burning and anthropogenic
352 emissions from southern Africa could have been transported to Rwanda after
353 photochemical aging and processing. During the DJF dry season, fires are closer to
354 Rwanda and away from major urban areas. Fire haze air masses thus likely
355 underwent less photochemical processing before arriving at the RCO and were
356 exposed to less anthropogenic O₃ precursors. Increased stagnation or higher



357 temperature effects are unlikely to be driving this observation as wind speed is higher
358 during JJA and temperature is similar or lower compared to DJF. During June and July,
359 a loose correlation ($R=0.47$ and 0.45 , respectively) between O_3 mixing ratios and BC
360 concentrations was observed, while no correlations ($R=-0.04$, -0.15 , and 0.07) were
361 observed in December, January, and February.

362 The MOZAIC campaign in the late 1990s and early 2000s measured ambient O_3
363 mixing ratios at the Nairobi, Kampala, and Kigali airports and found Kigali, despite its
364 smaller size and lower vehicle count, to have the highest O_3 mixing ratios among them
365 (Sauvage et al., 2005). They measured a similar in magnitude increase in surface O_3
366 mixing ratios during the JJA season in Rwanda as our measurements at the RCO,
367 although DJF was not measured in their work. Brazzaville, Republic of the Congo did
368 report January and February O_3 measurements during the MOZAIC campaign. While
369 much further west than Rwanda, in Brazzaville O_3 mixing ratios also increased during
370 January and February, parallel to Rwanda, with monthly averages during January and
371 February 25 ppb greater than the minimum of <30 ppb in April (Sauvage et al., 2005).
372 This suggests influence from northern hemisphere biomass burning to O_3 mixing
373 ratios at Brazzaville. Ozone in JJA at Brazzaville was almost 30 ppb higher than in
374 January and February, however, so transport of air mass from the south and southern
375 Africa biomass burning had a greater influence on O_3 in the region than transport
376 from the north and biomass burning in central Africa. The 1992 SAFARI campaign
377 also measured O_3 in sub-Saharan Africa throughout all seasons, and measured a
378 seasonal ozone concentration peak during the JJA period for central and southern
379 Africa (Thompson et al., 1996). A separate, large peak for DJF was not as observable



380 in the SAFARI data (Thompson et al., 1996). SAFARI measurements took place prior
381 to 1993, meaning that significant development in sub-Saharan Africa could have taken
382 place between the SAFARI campaign and the MOZAIC campaign (1997-2003) that
383 could drive the increasing O₃ in DJF as well as JJA over a period of almost a decade.
384 The SAFARI campaign measured the total column O₃, not the ground-level O₃ mixing
385 ratios, so data are not directly comparable.

386 The continuous collection of BC, CO and O₃ data during the dry and rainy
387 seasons allowed examination of both transported and local pollution in both seasons.
388 Here we define local pollution as pollution originating within twelve hours transport
389 time under typical wind speed conditions (<150 km, including both Rwanda and the
390 border areas with DRC and Uganda). During Rwanda's rainy seasons, the continental
391 fire count is also at a minimum, reducing large-scale biomass burning influence. The
392 region's emissions were from small-scale agricultural burning, charcoal making,
393 cooking fires, brick production (located in the valley below the station and throughout
394 the region), vehicles, diesel and heavy fuel-oil power plants, and diesel generators.
395 These activities continued throughout the rainy season and dry season at similar
396 rates.

397 The baseline daily average BC concentration in the rainy season remained at
398 0.5-1 µg m⁻³ after 12 hour periods without rain, which could be considered as
399 contributions of small but numerous diffuse emission sources to daily BC
400 concentration in this region. These values are not negligible, especially at a rural
401 location with little vehicular traffic or industries. If all BC during the rainy seasons is
402 assumed to be local in origin (within one day of transport, as typically rain occurs



403 each day during the rainy season), and this level remained the same throughout the
404 year, yearly average contribution of local emissions to BC would vary between 18-
405 100% of the total measured BC concentration at RCO-M. The shoulder months of
406 September and February have been removed from this calculation as they have both
407 rain and biomass burning influence, but on a yearly scale, around or greater than 35%
408 of BC concentration measured at the station could originate from local (day transport)
409 emissions, on par with previous estimates of the contribution of savanna and forest
410 burning BC emissions versus other emission sources (Bond et al., 2013). While
411 transported savanna, woodland, and forest fire emissions likely have a huge effect on
412 Rwanda's air quality, targeting local emissions could bring a marked decrease in PM
413 exposure of the population. More than 50% of BC measured in industrialized or
414 densely populated areas of South Africa is modeled to be from anthropogenic
415 emissions (Kuik et al., 2015). As Rwanda develops, its percentage contribution of
416 anthropogenic BC could increase like South Africa's, if there is no regulation, on top of
417 Rwanda experiencing 6 months a year of increased black carbon due to fire emissions
418 (3 months more influence than South Africa experiences). In other words, it is
419 necessary to reduce emissions from local sources within Rwanda to reduce exposure
420 to air pollution but this will not be sufficient to reduce exposure to WHO
421 recommended levels because Rwanda is significantly affected by large-scale regional
422 biomass burning pollution as well, most occurring outside of the country.

423 **3.2.1 Diurnal Variations in BC, CO and O₃**

424 Diurnal variations in pollutant concentration of pollutants can provide
425 important insights into information on local as well as regional pollution emission



426 sources. Diurnal variations in BC concentrations, CO mixing ratios and O₃ mixing
427 ratios observed at RCO in different by seasons are shown in Figure 8. At the RCO, the
428 O₃ mixing ratio exhibited a diurnal cycle with a peak in concentration in the evenings,
429 steady levels through the night and a minimum during mid-day. The increase of O₃ in
430 the evening is likely from regionally formed polluted air masses transported to the
431 station within the boundary layer and measured at the station during the collapsing
432 nocturnal boundary layer. Similar diurnal cycles were found at other mountain
433 locations remote from urban centers (Zhang et al., 2015). This diurnal pattern persists
434 in all seasons (Figure 8) and occurred on daily time scales, meaning that local urban
435 centers such as Kigali and Kampala were likely the source regions for this observed
436 daily increase in ozone. However, the differences in diurnal minima and maxima were
437 highest in the June-August period, and lowest in the December-February period. This
438 difference may be due to the differences in biomass burning proximity (far in JJA,
439 closer in DJF) and primary wind direction (southerly versus northerly). Flatter ozone
440 profiles generally indicate that ozone was neither produced locally nor heavily
441 influenced by daily variations in boundary layer height; as large-scale biomass
442 burning occurs closer to the RCO in DJF, O₃ may be more well-mixed within the
443 polluted mixed layer and have less time to be transported to higher altitudes in
444 troposphere as in JJA. Wind speed was similar during all months, and RH was similar
445 for DJF and JJA.

446 BC had mid-morning and early evening peaks that coincided with both cooking
447 times and kerosene/generator use times. The evening peak was higher, likely due to
448 more use of generator and kerosene lanterns for lighting in the evening (dark at 6



449 pm) than the morning (light at 6 am). These peaks persisted throughout the rainy and
450 dry seasons, indicating influence of local sources for these diurnal peaks. In the rainy
451 seasons, the average difference between minimum and maximum BC concentration
452 was around $0.5 \mu\text{g m}^{-3}$ and slightly higher ($0.6 \mu\text{g m}^{-3}$) in the dry season. During June-
453 August, we observed a general increase in BC concentrations in the morning hours,
454 while December-February, the morning had lower BC concentrations and the main
455 peak was observed in the evening with a small second peak in mid-afternoon. CO
456 mixing ratios had a similar but less pronounced diurnal variation. Like with O_3 ,
457 changing boundary layer conditions also likely played a role in variations in BC
458 concentrations over the day, as local boundary layer height increased during the day
459 and decreased during the evening and morning hours, and the RCO altitude was above
460 the boundary layer height often during the evening.

461 **3.3 BC source apportionment**

462 It is important to understand the pollution emission sources in East Africa,
463 beyond large-scale biomass burning, in order to enact policies and actions to reduce
464 these emissions. Past observational as well as atmospheric simulation studies show
465 that BC from biomass burning is the major source of pollution in this region (Andela
466 and van der Werf, 2014; Bond et al., 2013; Hersey et al., 2015; Marais and
467 Wiedinmyer, 2016; Sauvage et al., 2005), though fossil fuel BC emissions are
468 increasing with increasing socio-economic development activities (Knippertz et al.,
469 2015; Liousse et al., 2014). Regional emission models predict that only 10% of fuel
470 (by mass) burned in East Africa is fossil fuel (Marais and Wiedinmyer, 2016). Fuel
471 demand information from the Rwanda Bureau of Statistics and Ministry of



472 Infrastructure confirm <10% fossil fuel use is a reasonable amount for Rwanda, with
473 demand of fuel wood for both charcoal making and direct burning is 4.2 billion kg and
474 petroleum products is 287 million kg per year (wood use from 2016, fossil fuel use
475 extrapolated from statistics in 2011/12 using a 10% per annum growth rate, fuel use
476 summarized in Table 2)(Rwanda Ministry of Infrastructure, 2015). In regards to
477 demand of petroleum products, diesel is the highest (46%) followed by petrol (31%),
478 heavy fuel oil (12%), kerosene (6%), and jet-A fuel (5%). As of November 2016,
479 approximately 27% of Rwanda's on-grid power generation (52 MW out of 190 MW)
480 was either heavy fuel oil or diesel-derived ([http://www.reg.rw/index.php/our-](http://www.reg.rw/index.php/our-business/generation/624-power-generation-2)
481 [business/generation/624-power-generation-2](http://www.reg.rw/index.php/our-business/generation/624-power-generation-2)), the rest is a mix of hydroelectric
482 power (44%), methane gas (16%), solar, (5%) and imports from nearby countries.

483 While the majority of fuel use is wood, BC emission factors per kg fuel burned
484 are different for open burning, three stone fires, and diesel generators and other
485 engines using unregulated fuel in inefficient burning conditions (average
486 approximations of EF is 0.75, 1, and 4 g kg⁻¹, respectively)(FAO, 2010). As there are
487 less than 180,000 registered motor vehicles (including truck and motorcycles) in
488 Rwanda, and less than a third of the power generated in-country is from fossil fuel,
489 but over 2 million households use wood fuel for cooking, understanding the relative
490 contribution of each emission source to black carbon particulates is essential to guide
491 policy and development in the short-term. A four-times greater emission factor per
492 unit of fuel burned, even for fuel that only comprises 10% of the total fuel use in-
493 country, could make motor vehicle and diesel generator emissions a good first target.
494 However, burning conditions, fuel conditions, and engine combustion conditions



495 (including presence or absence of a diesel particle filter, age of the engine) can greatly
496 influence these emission factors. One way separation of fossil fuel combustion versus
497 biomass burning BC particulate has been estimated in the past is by measuring the
498 color of the particles (wood smoke particles have enhanced absorption in the UV,
499 while fossil fuel combustion particles have flat absorption over all
500 wavelengths)(Kirchstetter and Thatcher, 2012; Sandradewi et al., 2008).

501 The Aethalometer's seven wavelengths allow measurement of the wavelength-
502 dependent aerosol absorption coefficients that can be used to infer the potential
503 sources of BC aerosol (Drinovec et al., 2015; Sandradewi et al., 2008) measured.
504 Theoretically, from the wavelength dependence of aerosol absorption, BC from fossil
505 fuel and wood smoke can be differentiated(Sandradewi et al., 2008). Though this two-
506 component model can provide a valuable knowledge on knowledge on source
507 attribution of BC this model has some limitations. This model is more accurate if
508 calibrated to local conditions(Dumka et al., 2013; Harrison et al., 2012), as different
509 fuels and wood biomass burning creates aerosol with different radiative properties
510 and the standard model, based on European studies, has been shown to be less
511 applicable in developing countries (Garg et al., 2016).

512 From the Aethalometer data, wavelength dependence of absorption
513 coefficients and the absorption Ångstrom exponent (AAE) were calculated and
514 compared to literature values of different type of biomass burning and fossil fuel
515 combustion. The AAE is a dimensionless property commonly used to characterize the
516 wavelength-dependent absorption of BC and gives clues on the source and/or aging of
517 BC when compared to laboratory and other ambient studies(Chung et al., 2012; Lack



518 and Langridge, 2013; Russell et al., 2010; Yuan et al., 2016). The AAE values assigned
519 for the standard Aethalometer model separating the BC from biomass burning and
520 fossil fuel combustion are two and one, respectively (Kirchstetter et al, 2004;
521 Sandradewi et al, 2012; Drinovec et al. 2015). In this work, standard mass absorption
522 cross-sections (MACs) for each wavelength provided by the manufacturer of the
523 Aethalometer were used to calculate the absorption coefficient (b_{abs}) at each
524 wavelength. For pure BC from fossil fuel, $b_{\text{abs}} \sim 1/\lambda$ and the AAE between two
525 wavelengths (470 nm and 950 nm) is 1 using the equation $\ln(b_{\text{abs}}\lambda_1/b_{\text{abs}}\lambda_2)/\ln(\lambda_2/\lambda_1)$.
526 The average AAE (averaged for entire measurement period between July 2015 and
527 January 2017) was calculated to be 1.65 (+/- 0.14) at the RCO using the 470 and 950
528 wavelength absorption and MACs (Figure 9)(Sandradewi et al., 2008; Drinovec et al.
529 2015). These wavelengths were chosen as the AAE calculated from 470 and 950 is
530 generally comparable with other literature values(Saarikoski et al., 2012). The
531 calculated AAE values were on par with AAE calculated from measurements taken in
532 areas heavily influenced by biomass burning (Chung et al., 2012; Lack and Langridge,
533 2013; Russell et al., 2010; Saleh et al., 2013; Sandradewi et al., 2008; Yuan et al.,
534 2016). Past studies have reported an AAE of 1.2-2.5 for biomass burning
535 aerosol(Andreae and Gelencsér, 2006; Chung et al., 2012; Russell et al., 2010; Saleh et
536 al., 2013, 2014). While daily only small variations (+/- 0.05) for AAE were observed,
537 significant seasonal differences in this value were found, with monthly averaged
538 values ranging from 1.5 (dry season) to 1.9 (at the end of the long rainy season).
539 Studies in southern Africa measuring savanna and crop burning found an AAE of
540 around 1.45 for ambient black carbon aerosol, and in the dry season savanna and crop



541 burning are the prevalent type of large-scale biomass burning in sub-Saharan Africa
542 (Russell et al., 2010). The AAE calculated from the Aethalometer data at the RCO was
543 higher during the rainy season when local emissions dominated our measurements.
544 Eucalyptus burning, the most prevalent burning near the station (for charcoal making,
545 cooking fires, brick kiln fuel) was measured in laboratory experiments to have a
546 higher AAE than savanna burning (AAE of 1.71 ± 0.50 calculated between 405 and
547 781 nm wavelengths)(Chung et al., 2012). Eucalyptus trees and savanna burning were
548 certainly not the only two types of solid biofuel influencing measurements at the
549 station, but the difference in AAE of aerosols produced from different fuels means that
550 the AAE will have large variations based on fuel wood or other biomass used and this
551 was reflected in our data.

552 Using the Aethalometer model with standard inputs not accounting for the
553 different types of fuel used in East Africa versus Europe, a high influence of fossil fuel
554 black carbon emissions was calculated: in the dry season, over 50% of black carbon
555 was assigned to be fossil fuel in origin. Fossil fuel emissions certainly influenced the
556 pollution at the RCO, as air masses from Kigali, Kampala, Nairobi, and Dar es Salaam
557 were likely transported to the station and these cities likely have high black carbon
558 emissions from generators, fossil fuel power stations, and older diesel vehicles but
559 would also have significant biomass cook stove emissions (Gatari and Boman, 2003;
560 Koch et al., 2009; Mkoma et al., 2009; van Vliet and Kinney, 2007). However, this
561 number is likely falsely high: at <10% fuel demand of fossil fuel (all types) versus
562 >90% wood and charcoal fuel demand, even if the g BC per kg fuel from diesel was 4x
563 higher, and all fossil fuel use was unregulated diesel (unlikely), well under half of the



564 measured BC would be from fossil fuel combustion emissions. Additionally, the AAE
565 calculated in all seasons had a flat diurnal profile, suggesting no daily change in
566 aerosol source (e.g., rush hour traffic from Kigali) was measured at the RCO.

567 In order to gain more insights into the sources of BC we also examined the
568 BC:CO over time. CO is also released by inefficient combustion and the $\Delta\text{BC}:\Delta\text{CO}$ ratio
569 can be different for different emission sources. In order to calculate this ratio we first
570 converted the CO mixing ratios to concentrations (in $\mu\text{g m}^{-3}$), and then subtracted the
571 95th percentile values for CO and BC from their respective concentrations. For the
572 entire data set, the $\Delta\text{BC}:\Delta\text{CO}$ (both in $\mu\text{g m}^{-3}$) ratio was 0.014 (R^2 0.79, $n = 40523$).
573 The $\Delta\text{BC}:\Delta\text{CO}$ ratio varied seasonally, with monthly average peaks reaching 0.016 in
574 December, February, and July and lows below 0.01 in April. The average ratio of 0.014
575 for the measurement period was almost twice as high as in biomass burning plumes
576 sampled over West Africa in an aircraft campaign (0.0072)(Moosmüller and
577 Chakrabarty, 2011) but on par with or lower than measurements taken during the
578 INDOEX campaign in the Indian Ocean (Dickerson et al., 2002). A study in Germany
579 and Mexico found a correlation between diesel vehicle use and higher BC:CO
580 (Baumgardner et al., 2002), while other studies have also found an increased
581 $\Delta\text{BC}:\Delta\text{CO}$ during periods more influenced by biomass burning (Pan et al., 2011). A
582 study in India found no correlation in biomass-burning and fossil fuel-influenced
583 $\Delta\text{BC}:\Delta\text{CO}$ air masses (Sahu et al., 2012), as there are a wide range of ratios measured
584 from the same source (Dickerson et al., 2002; Sahu et al., 2012). The high $\Delta\text{BC}:\Delta\text{CO}$
585 ratio at the RCO could be due to the prevalence of older diesel engines in the country,
586 which emit more BC to CO than newer engines (Cai et al., 2013), but, as the highest



587 value occurs during the Rwanda dry seasons and the continental biomass burning
588 seasons, likely the ratio is governed in part by rainout as BC is more easily removed
589 by wet deposition than CO. In this study, we were not able to use this ratio to further
590 separate biomass burning BC from fossil fuel combustion BC, likely due to differences
591 in emission profiles in developing versus developed countries. This underscores the
592 need for more measurements in East Africa to understand emission sources and
593 profiles and develop more-robust emission source profiles.

594 **4. Conclusions**

595 In this work, we present the first long-term and highly time-resolved
596 continuous measurements of short-lived climate forcers for a nearly two-year period
597 from July 2015 to January 2017 at the Rwanda Climate Observatory located at Mt.
598 Mugogo in Rwanda, in one of the data poor regions of the world. From these
599 observations, we find that:

- 600 1. During Rwanda's two dry seasons, the country experienced pollution
601 transported from both the northern (DJF) and southern (JJA) biomass
602 burning seasons in Africa. This transported pollution led to high black
603 carbon and carbon monoxide levels at the RCO, surpassing
604 concentrations measured in many major cities elsewhere. Emissions
605 from large-scale crop and savanna fires appeared to have a wide-
606 reaching effect on this region, reflected in increased PM_{2.5} in
607 Kampala, a major East African city, for both biomass burning seasons
608 and likely driving the increased O₃ measured during DJF and JJA by
609 our study and by past studies in equatorial Africa. The dense



610 population of equatorial East Africa and the double impact of the two
611 fires seasons could lead to significant public health problems for the
612 population in Rwanda and equatorial East Africa as exposure to
613 elevated levels of PM_{2.5} and BC concentrations occurs six months out
614 of the year.

615 2. Ground level O₃ was enhanced during both dry seasons, likely due to
616 the prevalent wide-scale biomass burning. Increased enhancement
617 was observed during the JJA dry season when the air masses
618 originated from the southeast and likely included a mix of biomass
619 burning and anthropogenic emissions (cooking fires, vehicles,
620 industries), leading to higher ozone concentrations at the RCO
621 downwind of these mixed pollution sources. As this area develops
622 and population grows, local as well as regional air pollution could
623 become a major environmental and societal issue that could be a
624 threat to national development goals.

625 3. Local emissions beyond large-scale biomass burning influence were
626 constant and estimated to contribute between 18-almost 100 % of
627 the measured black carbon concentrations depending on season, if
628 black carbon during the rainy season was assumed to be completely
629 local (Rwanda and neighboring countries) in origin (ranging from
630 0.5-1 µg m⁻³ daily average measured BC). These local emissions, from
631 different combustion sources (e.g., cooking fires, inefficient diesel
632 generators and engines with sub-standard fuel use, solid biomass fuel



633 burning, small agricultural fires), are likely concentrated in the
634 densely populated Rwanda and Lake Kivu economic area. Rwanda's
635 population is growing quickly and, as these local emissions are
636 related to population density, air pollution will likely increase unless
637 there is government intervention.

638 4. Different combustion fuel and burning practices in Europe and East
639 Africa calls into question the accuracy and applicability of a two-
640 component model for estimating BC from fossil fuel combustion and
641 biomass burning using AAE approximations for biomass burning and
642 fossil fuel combustion aerosol measured in Europe for use in East
643 Africa. There may also be different mass absorption cross-sections
644 for aerosols measured at the RCO than in Europe or North America.
645 This shows the need for multiple on-ground measurements to fully
646 understand pollution sources in different regions of the world,
647 notably in Africa. However, seasonal variations in the wavelength
648 dependence of ambient BC particles did point to different sources of
649 BC particles and this should be further explored in future studies.

650 5. The measurements we have provided in this study will be useful in
651 advancing atmospheric science in Africa, improve emission
652 inventories and air pollution/atmospheric models in the region, and
653 designing mitigation measures in the region, which has limited long-
654 term and in-situ atmospheric data.

655



656 These data and analyses, while acknowledging the high influence of regional
657 biomass burning, also show that significant decreases in air pollution could be
658 achieved within eastern and central Africa with targeted local policies, emphasizing
659 cleaner diesel vehicles and generators, reduced wood-fuel reliance for cook stoves,
660 and improved cook stoves to burn biomass fuel more efficiently. Currently, over 2
661 million households in Rwanda rely on wood burning (including charcoal) for cooking.
662 While reducing this number will have significant economic costs, putting in place
663 infrastructure for alternative cooking fuels (pellet stoves, LPG stoves, electrical
664 stoves) could help the country avoid even higher local air pollution emissions and
665 associated adverse impacts as the population grows. Diesel-fueled minibuses,
666 common transport between towns in Rwanda and within Kigali, and older diesel
667 vehicles are also high emitters of black carbon but newer vehicles with emissions
668 control technology may be economically beyond the reach of local bus companies and
669 citizens. Continuing to grow electrical capacity and connection will reduce the use of
670 kerosene lanterns and diesel generators, and will reduce air pollution if additional
671 energy capacity is achieved through renewable sources (solar, hydropower). The
672 huge influence of regional biomass burning, exacerbated by equatorial East Africa's
673 meteorology, and the potential influence of anthropogenic emissions from major
674 cities on O₃ formation in this regions must also be examined as this area develops and
675 this should be an important agenda for the regional discussions on environmental,
676 public health, and other development issues.

677

678 **6. Future Work**



679 The government of Rwanda is working to establish an air quality and climate
680 change monitoring network throughout the country to measure ambient criteria air
681 pollutants and other key climate change related components of atmospheric pollution.
682 Building knowledge of air quality and climate change related emissions in this data-
683 poor area of the world is essential to fill the large data and knowledge gap in this
684 region. Adding ground-based measurements, comparing measurements to satellite
685 data, using data to evaluate and improve existing emission inventories, improving
686 accuracy of global/regional air quality and climate change models, and using data for
687 quantification of impacts of air pollution and climate change will help local
688 governments design appropriate mitigation strategies rooted in data and local
689 context.

690 **7. Data Availability**

691 This data will be made available at the AGAGE website,
692 <https://agage.mit.edu/data/agage-data>. All data used in this article will be made
693 available as of publication and data from this project on a rolling basis after quality
694 control.

695 **Acknowledgments:**

696 We thank the generous MIT alumni donors to the MIT-Rwanda Climate Observatory
697 Project that provided the funds to purchase, develop and install most of the
698 instruments at the Rwanda Climate Observatory. Additional funds for this purpose
699 were provided by the MIT Center for Global Change Science. COMESA provided the
700 funds to purchase and install the Aethalometer at the RCO. We also thank the
701 Government of Rwanda and the Rwanda Ministry of Education, specifically Mike



702 Hughes, Vianney Rugamba, and Dr. Marie Christine Gasingirwa, for supporting this
703 project, including funding the staffing and infrastructure costs of the Rwanda Climate
704 Observatory and the University of Rwanda for providing laboratory space and
705 infrastructure for instrument testing. We thank Dr. Arnico Panday who provided
706 guidance during the initial stages of this project. We also wish to acknowledge the
707 essential contributions of the Mugogo station technical experts Theobard Habineza,
708 Modeste Mugabo, Olivier Shyaka, and Gaston Munyampundu, and RBA technician
709 Yves Fidele, without which running this station would be impossible.

710

711



712 Table 1: Instruments used in this study and measurement period used for analysis

INSTRUMENT	SPECIES MEASURED	MEASUREMENT PERIOD	TIME RESOLUTION
PICARRO G2401 CAVITY RING DOWN SPECTROMETER	CO ₂ , CO, CH ₄ , H ₂ O	MAY 2015-JANUARY 2017	1 MIN
MAGEE SCIENTIFIC AE33 7-WAVELENGTH AETHALOMETER	BLACK CARBON (PM _{2.5} , CYCLONE IMPACTOR ON INLET)	MAY 2015-JANUARY 2017	1 MIN
TELEDYNE T400 API	O ₃	MAY 2015-JANUARY 2017	1 MIN
VAISALA WXT	MET PARAMETERS (RH, WS, WD, T, P)	JULY 2015-JANUARY 2017	1S



713

714 Table 2:

715

716 Fuel Demand in Rwanda (2016, Rwanda Ministry of Infrastructure)

Fuel Type	Demand
Petrol	120442 kL
Diesel	178529 kL
Kerosene	22288 kL
Heavy Fuel Oils	59292 kL
Jet-A	18235 kL
Wood (charcoal + natural)	4,200,000 metric tons

717

718



719 References

720

721 Andela, N. and van der Werf, G. R.: Recent trends in African fires driven by cropland
722 expansion and El Niño to La Niña transition, *Nat. Clim. Chang.*, 4(9), 791–795,
723 doi:10.1038/nclimate2313, 2014.

724 Andreae, M. O. and Gelencsér, A.: Black carbon or brown carbon? The nature of light-
725 absorbing carbonaceous aerosols, *Atmos. Chem. Phys.*, 6(3), 3419–3463,
726 doi:10.5194/acpd-6-3419-2006, 2006.

727 Archibald, S., Nickless, A., Govender, N., RJ., S. and Lehsten, V.: Climate and the inter-
728 annual variability of fire in southern Africa: a meta-analysis using long-term field
729 data and satellite-derived burnt area data, *Glob. Ecol. Biogeogr.*, 19(6), 794–809,
730 2010.

731 Baumgardner, D., Raga, G., Peralta, O., Rosas, I., Castro, T., Kuhlbusch, T., John, A. and
732 Petzold, A.: Diagnosing black carbon trends in large urban areas using carbon
733 monoxide measurements, *J. Geophys. Res. Atmos.*, 107(21),
734 doi:10.1029/2001JD000626, 2002.

735 Bond, T. C., Doherty, S. J., Fahey, D. W., Forster, P. M., Berntsen, T., DeAngelo, B. J.,
736 Flanner, M. G., Ghan, S., Kärcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P. K.,
737 Sarofim, M. C., Schultz, M. G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S.,
738 Bellouin, N., Guttikunda, S. K., Hopke, P. K., Jacobson, M. Z., Kaiser, J. W., Klimont,
739 Z., Lohmann, U., Schwarz, J. P., Shindell, D., Storelvmo, T., Warren, S. G. and Zender,
740 C. S.: Bounding the role of black carbon in the climate system: A scientific
741 assessment, *J. Geophys. Res. Atmos.*, 118(11), 5380–5552,
742 doi:10.1002/jgrd.50171, 2013.

743 Cai, H., Burnham, A. and Wang, M.: Updated Emission Factors of Air Pollutants from
744 Vehicle Operations in GREET TM Using MOVES, , (September), 2013.

745 Carslaw, D. C. . and Ropkins, K.: The openair manual open-source tools for analysing
746 air pollution data, *King's Coll. London*, 27–28(January), 287, 2012.

747 Chung, C. E., Kim, S. W., Lee, M., Yoon, S. C. and Lee, S.: Carbonaceous aerosol AAE
748 inferred from in-situ aerosol measurements at the Gosan ABC super site, and the
749 implications for brown carbon aerosol, *Atmos. Chem. Phys.*, 12(14), 6173–6184,
750 doi:10.5194/acp-12-6173-2012, 2012.

751 Crutzen, P. J. and Andreae, M.: Biomass Burning in the Tropics : Impact on
752 Atmospheric Chemistry and Biogeochemical Cycles Estimates of Worldwide
753 Biomass Burning, *Science (80-.)*, 250(4988), 1669–1678,
754 doi:10.1126/science.250.4988.1669, 1990.

755 Dickerson, R. R., Andreae, M. O., Campos, T., Mayol-Bracero, O. L., Neusuess, C. and
756 Streets, D. G.: Analysis of black carbon and carbon monoxide observed over the
757 Indian Ocean: Implications for emissions and photochemistry, *J. Geophys. Res.*,
758 107(D19), doi:Artn 8017\rDoi 10.1029/2001jd000501, 2002.

759 Drinovec, L., Močnik, G., Zotter, P., Prévôt, A. S. H., Ruckstuhl, C., Coz, E., Rupakheti, M.,
760 Sciare, J., Müller, T., Wiedensohler, A. and Hansen, A. D. A.: The “dual-spot”
761 Aethalometer: An improved measurement of aerosol black carbon with real-time
762 loading compensation, *Atmos. Meas. Tech.*, 8(5), 1965–1979, doi:10.5194/amt-8-
763 1965-2015, 2015.

764 Dumka, U. C., Manchanda, R. K., Sinha, P. R., Sreenivasan, S., Moorthy, K. K. and Suresh



- 765 Babu, S.: Temporal variability and radiative impact of black carbon aerosol over
766 tropical urban station Hyderabad, *J. Atmos. Solar-Terrestrial Phys.*, 105–
767 106(April 2016), 81–90, doi:10.1016/j.jastp.2013.08.003, 2013.
- 768 Field, R. D., van der Werf, G. R., Fanin, T., Fetzner, E. J., Fuller, R., Jethva, H., Levy, R.,
769 Livesey, N. J., Luo, M., Torres, O. and Worden, H. M.: Indonesian fire activity and
770 smoke pollution in 2015 show persistent nonlinear sensitivity to El Niño-induced
771 drought, *Proc. Natl. Acad. Sci.*, 113(33), 9204–9209,
772 doi:10.1073/pnas.1524888113, 2016.
- 773 Food and Agriculture Organization of the United Nations: What woodfuels can do to
774 mitigate climate change, *FAO For. Pap.*, 98, 2010.
- 775 Galbraith, K.: Measuring Africa's Air Pollution, *New York Times*, 16th April [online]
776 Available from: [http://www.nytimes.com/2014/04/17/business/energy-
777 environment/measuring-africas-air-pollution.html?_r=0](http://www.nytimes.com/2014/04/17/business/energy-environment/measuring-africas-air-pollution.html?_r=0), 2014.
- 778 Garg, S., Chandra, B. P., Sinha, V., Sarda-Esteve, R., Gros, V. and Sinha, B.: Limitation of
779 the Use of the Absorption Angstrom Exponent for Source Apportionment of
780 Equivalent Black Carbon: a Case Study from the North West Indo-Gangetic Plain,
781 *Environ. Sci. Technol.*, 50(2), 814–824, doi:10.1021/acs.est.5b03868, 2016.
- 782 Gasore, J.: Quantifying Emissions of Carbon Dioxide and Methane in Central and
783 Eastern Africa Through High Frequency Measurements and Inverse Modeling,
784 *Massachusetts Institute of Technology.*, 2018.
- 785 Gatari, M. J. and Boman, J.: Black carbon and total carbon measurements at urban and
786 rural sites in Kenya, East Africa, *Atmos. Environ.*, 37(8), 1149–1154,
787 doi:10.1016/S1352-2310(02)01001-4, 2003.
- 788 Harrison, R. M., Beddows, D. C. S., Hu, L. and Yin, J.: Comparison of methods for
789 evaluation of wood smoke and estimation of UK ambient concentrations, *Atmos.*
790 *Chem. Phys.*, 12(17), 8271–8283, doi:10.5194/acp-12-8271-2012, 2012.
- 791 Hersey, S. P., Garland, R. M., Crosbie, E., Shingler, T., Sorooshian, A., Piketh, S. and
792 Burger, R.: An overview of regional and local characteristics of aerosols in South
793 Africa using satellite, ground, and modeling data, *Atmos. Chem. Phys.*, 15(8),
794 4259–4278, doi:10.5194/acp-15-4259-2015, 2015.
- 795 Hsu, Y. K., Holsen, T. M. and Hopke, P. K.: Comparison of hybrid receptor models to
796 locate PCB sources in Chicago, *Atmos. Environ.*, 37(4), 545–562,
797 doi:10.1016/S1352-2310(02)00886-5, 2003.
- 798 IHME: GBD Compare Data Visualization, *Inst. Heal. Metrics Eval. Seattle, WA IHME,*
799 *Univ. Washingt.*, 2016.
- 800 Kabashnikov, V. P., Chaikovskiy, A. P., Kucsera, T. L. and Metelskaya, N. S.: Estimated
801 accuracy of three common trajectory statistical methods, *Atmos. Environ.*, 45(31),
802 5425–5430, doi:10.1016/j.atmosenv.2011.07.006, 2011.
- 803 Kalnay, E., Kanamitsu, M., Kistler, R., Collins, W., Deaven, D., Gandin, L., Iredell, M.,
804 Saha, S., White, G., Woollen, J., Zhu, Y., Chelliah, M., Ebisuzaki, W., Higgins, W.,
805 Janowiak, J., Mo, K. C., Ropelewski, C., Wang, J., Leetmaa, A., Reynolds, R., Jenne, R.
806 and Joseph, D.: The NCEP/NCAR 40-year reanalysis project, *Bull. Am. Meteorol.*
807 *Soc.*, 77(3), 437–471, doi:10.1175/1520-0477(1996)077<0437:TNYRP>2.0.CO;2,
808 1996.
- 809 Kinney, P. L., Gichuru, M. G., Volavka-Close, N., Ngo, N., Ndiba, P. K., Law, A., Gachanja,
810 A., Gaita, S. M., Chillrud, S. N. and Sclar, E.: Traffic impacts on PM_{2.5} air quality in



- 811 Nairobi, Kenya, *Environ. Sci. Policy*, 14(4), 369–378,
812 doi:10.1016/j.envsci.2011.02.005, 2011.
- 813 Kirchstetter, T. W. and Thatcher, T. L.: Contribution of organic carbon to wood smoke
814 particulate matter absorption of solar radiation, *Atmos. Chem. Phys.*, 12(14),
815 6067–6072, doi:10.5194/acp-12-6067-2012, 2012.
- 816 Knippertz, P., Coe, H., Chiu, J. C., Evans, M. J., Fink, A. H., Kalthoff, N., Liousse, C., Mari,
817 C., Allan, R. P., Brooks, B., Danour, S., Flamant, C., Jegede, O. O., Lohou, F. and
818 Marsham, J. H.: The daccwa project: Dynamics-aerosol-chemistry-cloud
819 interactions in West Africa, *Bull. Am. Meteorol. Soc.*, 96(9), 1451–1460,
820 doi:10.1175/BAMS-D-14-00108.1, 2015.
- 821 Koch, D., Schulz, M., Kinne, S., McNaughton, C., Spackman, J. R., Balkanski, Y., Bauer, S.,
822 Berntsen, T., Bond, T. C., Boucher, O., Chin, M., Clarke, A., De Luca, N., Dentener, F.,
823 Diehl, T., Dubovik, O., Easter, R., Fahey, D. W., Feichter, J., Fillmore, D., Freitag, S.,
824 Ghan, S., Ginoux, P., Gong, S., Horowitz, L., Iversen, T., Kirkevåg, A., Klimont,
825 Z., Kondo, Y., Krol, M., Liu, X., Miller, R., Montanaro, V., Moteki, N., Myhre, G.,
826 Penner, J. E., Perlwitz, J., Pitari, G., Reddy, S., Sahu, L., Sakamoto, H., Schuster, G.,
827 Schwarz, J. P., Seland, Ø., Stier, P., Takegawa, N., Takemura, T., Textor, C., van
828 Aardenne, J. a. and Zhao, Y.: Evaluation of black carbon estimations in global
829 aerosol models, *Atmos. Chem. Phys.*, 9(22), 9001–9026, doi:10.5194/acp-9-9001-
830 2009, 2009.
- 831 Kuik, F., Lauer, A., Beukes, J. P., Zyl, P. G. Van, Josipovic, M., Vakkari, V., Laakso, L. and
832 Feig, G. T.: The anthropogenic contribution to atmospheric black carbon
833 concentrations in southern Africa : a WRF-Chem modeling study, , 8809–8830,
834 doi:10.5194/acp-15-8809-2015, 2015.
- 835 Lack, D. A. and Langridge, J. M.: On the attribution of black and brown carbon light
836 absorption using the aerosol angstrom exponent, *Atmos. Chem. Phys.*, 13(20),
837 10535–10543, doi:10.5194/acp-13-10535-2013, 2013.
- 838 Liousse, C., Assamoi, E., Criqui, P., Granier, C. and Rosset, R.: Explosive growth in
839 African combustion emissions from 2005 to 2030, *Environ. Res. Lett.*, 9(3),
840 35003, doi:10.1088/1748-9326/9/3/035003, 2014.
- 841 Lupu, A. and Maenhaut, W.: Application and comparison of two statistical trajectory
842 techniques for identification of source regions of atmospheric aerosol species,
843 *Atmos. Environ.*, 36(36–37), 5607–5618, doi:10.1016/S1352-2310(02)00697-0,
844 2002.
- 845 Marais, E. A. and Wiedinmyer, C.: Air Quality Impact of Diesel use and Inefficient
846 Combustion Emissions in Africa (DICE-Africa), , doi:10.1021/acs.est.6b02602,
847 2016.
- 848 Mkoma, S. L., Maenhaut, W., Chi, X., Wang, W. and Raes, N.: Characterisation of PM10
849 atmospheric aerosols for the wet season 2005 at two sites in East Africa, *Atmos.*
850 *Environ.*, 43(3), 631–639, doi:10.1016/j.atmosenv.2008.10.008, 2009.
- 851 Moosmüller, H. and Chakrabarty, R. K.: Technical Note: Simple analytical relationships
852 between Ångström coefficients of aerosol extinction, scattering, absorption, and
853 single scattering albedo, *Atmos. Chem. Phys.*, 11(20), 10677–10680,
854 doi:10.5194/acp-11-10677-2011, 2011.
- 855 Ngo, N. S., Gatari, M., Yan, B., Chillrud, S. N., Bouhamam, K. and Kinney, P. L.:
856 Occupational exposure to roadway emissions and inside informal settlements in



- 857 sub-Saharan Africa: A pilot study in Nairobi, Kenya, *Atmos. Environ.*, 111, 179–
858 184, doi:10.1016/j.atmosenv.2015.04.008, 2015.
- 859 Niang, I., Ruppel, O. C., Abdrabo, M. A., Essel, A., Lennard, C., Padgham, J. and Urquhart,
860 P.: Africa, *Clim. Chang. 2014 Impacts, Adapt. Vulnerability - Contrib. Work. Gr. II*
861 *to Fifth Assess. Rep. Intergov. Panel Clim. Chang.*, 1199–1265,
862 doi:10.1017/CBO9781107415386.002, 2014.
- 863 Pan, X. L., Kanaya, Y., Wang, Z. F., Liu, Y., Pochanart, P., Akimoto, H., Sun, Y. L., Dong, H.
864 B., Li, J., Irie, H. and Takigawa, M.: Correlation of black carbon aerosol and carbon
865 monoxide in the high-altitude environment of Mt. Huang in Eastern China, *Atmos.*
866 *Chem. Phys.*, 11(18), 9735–9747, doi:10.5194/acp-11-9735-2011, 2011.
- 867 Prinn, R. G., Weiss, R. F., Fraser, P. J., Simmonds, P. G., Cunnold, D. M., Aleya, F. N.,
868 O'Doherty, S., Salameh, P., Miller, B. R., Huang, J., Wang, R. H. J., Hartley, D. E.,
869 Harth, C., Steele, L. P., Sturrock, G., Midgley, P. M. and McCulloch, A.: A history of
870 chemically and radiatively important gases in air deduced from
871 ALE/GAGE/AGAGE, *J. Geophys. Res. Atmos.*, 105(D14), 17751–17792,
872 doi:10.1029/2000JD900141, 2000.
- 873 Ramanathan, V. and Carmichael, G.: Global and regional climate changes due to black
874 carbon, *Nat. Geosci.*, 1, 221–227, doi:10.1038/ngeo156, 2008.
- 875 Real, E., Orlandi, E., Law, K. S., Fierli, F., Josset, D., Cairo, F., Schlager, H., Borrmann, S.,
876 Kunkel, D., Volk, C. M., McQuaid, J. B., Stewart, D. J., Lee, J., Lewis, A. C., Hopkins, J.
877 R., Ravegnani, F., Ulanovski, A. and Liousse, C.: Cross-hemispheric transport of
878 central African biomass burning pollutants: Implications for downwind ozone
879 production, *Atmos. Chem. Phys.*, 10(6), 3027–3046, doi:10.5194/acpd-9-17385-
880 2009, 2010.
- 881 Russell, P. B., Bergstrom, R. W., Shinozuka, Y., Clarke, a. D., DeCarlo, P. F., Jimenez, J. L.,
882 Livingston, J. M., Redemann, J., Holben, B., Dubovik, O. and Strawa, A.: Absorption
883 Angstrom Exponent in AERONET and related data as an indicator of aerosol
884 composition, *Atmos. Chem. Phys.*, 10, 1156–1169, doi:10.5194/acpd-9-21785-
885 2009, 2010.
- 886 Rwanda, R. of: Energy Sector Strategic Plan: Republic of Rwanda Ministry of
887 Infrastructure, Kigali, Rwanda., 2015.
- 888 Saarikoski, S., Carbone, S., Decesari, S., Giulianelli, L., Angelini, F., Canagaratna, M., Ng,
889 N. L., Trimborn, a., Facchini, M. C., Fuzzi, S., Hillamo, R. and Worsnop, D.: Chemical
890 characterization of springtime submicrometer aerosol in Po Valley, Italy, *Atmos.*
891 *Chem. Phys.*, 12(18), 8401–8421, doi:10.5194/acp-12-8401-2012, 2012.
- 892 Sahu, L. K., Kondo, Y., Moteki, N., Takegawa, N., Zhao, Y., Cubison, M. J., Jimenez, J. L.,
893 Vay, S., Diskin, G. S., Wisthaler, A., Mikoviny, T., Huey, L. G., Weinheimer, A. J. and
894 Knapp, D. J.: Emission characteristics of black carbon in anthropogenic and
895 biomass burning plumes over California during ARCTAS-CARB 2008, *J. Geophys.*
896 *Res. Atmos.*, 117(16), 1–20, doi:10.1029/2011JD017401, 2012.
- 897 Saleh, R., Hennigan, C. J., McMeeking, G. R., Chuang, W. K., Robinson, E. S., Coe, H.,
898 Donahue, N. M. and Robinson, A. L.: Absorptivity of brown carbon in fresh and
899 photo-chemically aged biomass-burning emissions, *Atmos. Chem. Phys.*, 13(15),
900 7683–7693, doi:10.5194/acp-13-7683-2013, 2013.
- 901 Saleh, R., Robinson, E. S., Tkacik, D. S., Ahern, A. T., Liu, S., Aiken, A. C., Sullivan, R. C.,
902 Presto, A. a., Dubey, M. K., Yokelson, R. J., Donahue, N. M. and Robinson, A. L.:



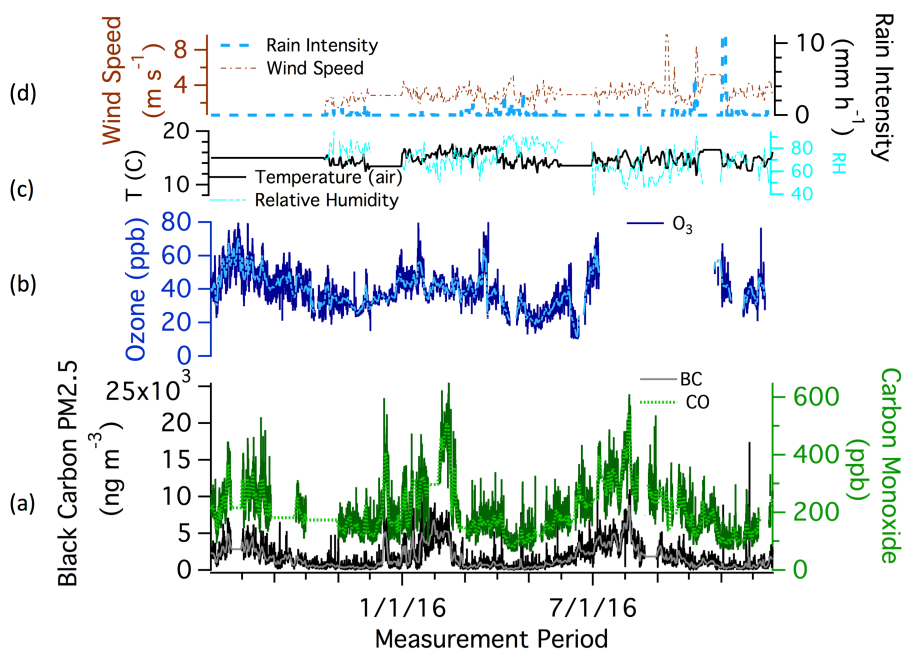
- 903 Brownness of organics in aerosols from biomass burning linked to their black
904 carbon content, *Nat. Geosci.*, 7(September), 1–4, doi:10.1038/ngeo2220, 2014.
- 905 Sandradewi, J., Prévôt, A. S. H., Szidat, S., Perron, N., Alfarra, M. R., Lanz, V. A.,
906 Weingartner, E. and Baltensperger, U.: Using Aerosol Light Absorption
907 Measurements for the Quantitative Determination of Wood Burning and Traffic
908 Emission Contributions to Particulate Matter, *Environ. Sci. Technol.*, 42(9), 3316–
909 3323, doi:10.1021/es702253m, 2008.
- 910 Sauvage, B., Thouret, V., Cammas, J. P., Gheusi, F., Athier, G. and Nédélec, P.:
911 Tropospheric ozone over Equatorial Africa: regional aspects from the MOZAIC
912 data, *Atmos. Chem. Phys.*, 5, 311–335, doi:10.5194/acpd-4-3285-2004, 2005.
- 913 Seibert, P., Kromp-Kolb, H., Baltensperger, U., Jost, D. T. and Schwikowski, M.:
914 Trajectory Analysis of High-Alpine Air Pollution Data, in *Air Pollution Modeling
915 and Its Application: NATO: Challenges of Modern Society*, edited by S.-E. (Riso N.
916 L. Gryning and M. M. (Centre for E. S. of the M. Millan, pp. 595–596, Springer USA,
917 1994.
- 918 Thompson, A. M., Diab, R. D., Bodeker, G. E., Zunckel, M., Coetzee, G. J. R., Archer, C. B.,
919 Mcnamara, D. P., Pickering, K. E., Combrink, J., Fishman, J. and Nganga, D.: Ozone
920 over southern Africa during SAFARI-92 / TRACE A, , 101(95), 1996.
- 921 Tiitta, P., Vakkari, V., Croteau, P., Beukes, J. P., Van Zyl, P. G., Josipovic, M., Venter, A. D.,
922 Jaars, K., Pienaar, J. J., Ng, N. L., Canagaratna, M. R., Jayne, J. T., Kermit, V. M.,
923 Kokkola, H., Kulmala, M., Laaksonen, A., Worsnop, D. R. and Laakso, L.: Chemical
924 composition, main sources and temporal variability of PM1 aerosols in southern
925 African grassland, *Atmos. Chem. Phys.*, 14(4), 1909–1927, doi:10.5194/acp-14-
926 1909-2014, 2014.
- 927 van Vliet, E. D. S. and Kinney, P. L.: Impacts of roadway emissions on urban particulate
928 matter concentrations in sub-Saharan Africa: new evidence from Nairobi, Kenya,
929 *Environ. Res. Lett.*, 2(4), 45028, doi:10.1088/1748-9326/2/4/045028, 2007.
- 930 Welp, L. R., Keeling, R. F., Weiss, R. F., Paplawsky, W. and Heckman, S.: Design and
931 performance of a Nafion dryer for continuous operation at CO₂ and CH₄ air
932 monitoring sites, *Atmos. Meas. Tech.*, 6(5), 1217–1226, doi:10.5194/amt-6-1217-
933 2013, 2013.
- 934 WHO: Health Effects of Particulate Matter: Policy implications for countries in eastern
935 Europe, Caucasus and central Asia, *World Heal. Organ.*, 15 [online] Available
936 from: www.euro.who.int, 2013.
- 937 World Bank: World Development Report 2011: World Development Indicators, Fossil
938 Fuel Energy Consumption., 2011.
- 939 Yuan, J. F., Huang, X. F., Cao, L. M., Cui, J., Zhu, Q., Huang, C. N., Lan, Z. J. and He, L. Y.:
940 Light absorption of brown carbon aerosol in the PRD region of China, *Atmos.
941 Chem. Phys.*, 16(3), 1433–1443, doi:10.5194/acp-16-1433-2016, 2016.
- 942 Zhang, L., Jin, L., Zhao, T., Yin, Y., Zhu, B., Shan, Y., Guo, X., Tan, C., Gao, J. and Wang, H.:
943 Diurnal variation of surface ozone in mountainous areas: Case study of Mt. Huang,
944 East China, *Sci. Total Environ.*, 538, 583–590,
945 doi:10.1016/j.scitotenv.2015.08.096, 2015.
- 946 Zhang, Y., Cooper, O. R., Gaudel, A., Thompson, A. M., Nédélec, P., Ogino, S. and West, J.
947 J.: Tropospheric ozone change from 1980 to 2010 dominated by equatorward
948 redistribution of emissions, *Nat. Geosci.*, 9(December), 875–881,



949 doi:10.1038/NCEO2827, 2016.
950
951



952 **Figure 1.** From top left moving counter-clockwise: an aerial view of RCO at Mt.
953 Mugogo Main Peak, the station with towers in the background, and the location of Mt.
954 Mugogo in Rwanda (blue pin) in relation to Kigali (yellow pin).
955



956

957

958 **Figure 2.** From the bottom up: (a) black carbon (black, grey) and carbon monoxide
959 (dark green, light green) (15 minute, daily) average concentrations; (b) ozone (dark
960 blue, light blue) (15 minute, daily), (c) temperature (black) and relative humidity
961 (light blue) values; (d) wind speed (red dotted) and rain intensity (blue dash) daily
962 average values.

963

964

965

966

967

968

969

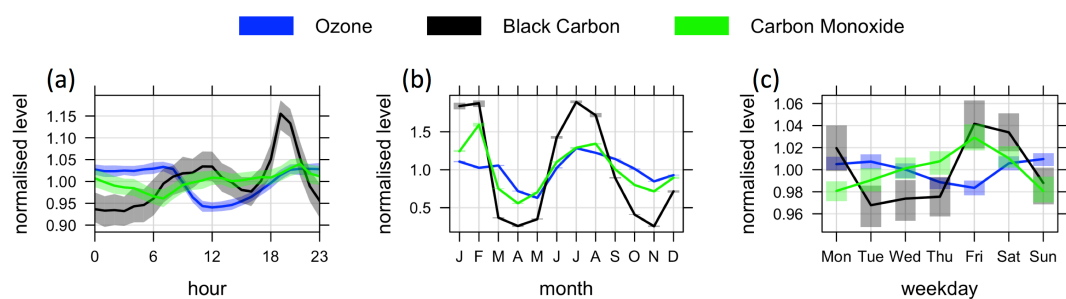
970

971

972

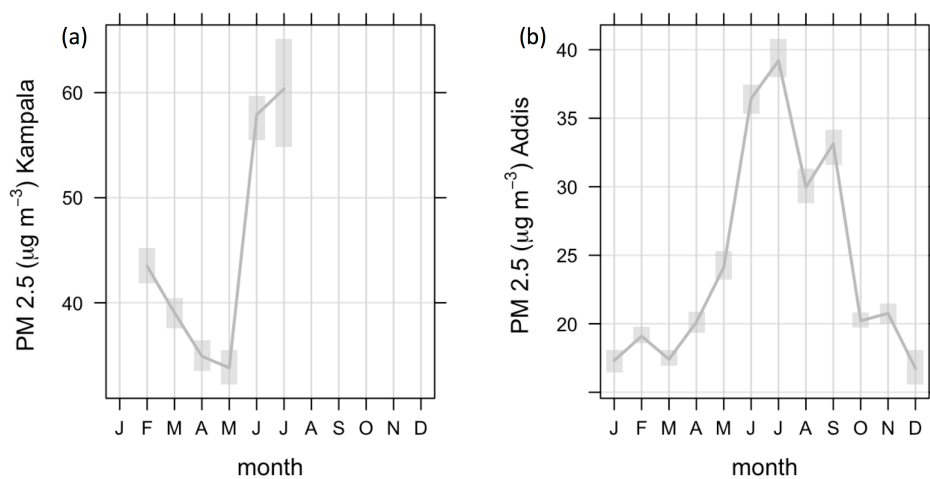


973
974
975
976
977
978
979
980
981



982
983
984
985
986
987
988
989
990
991
992
993

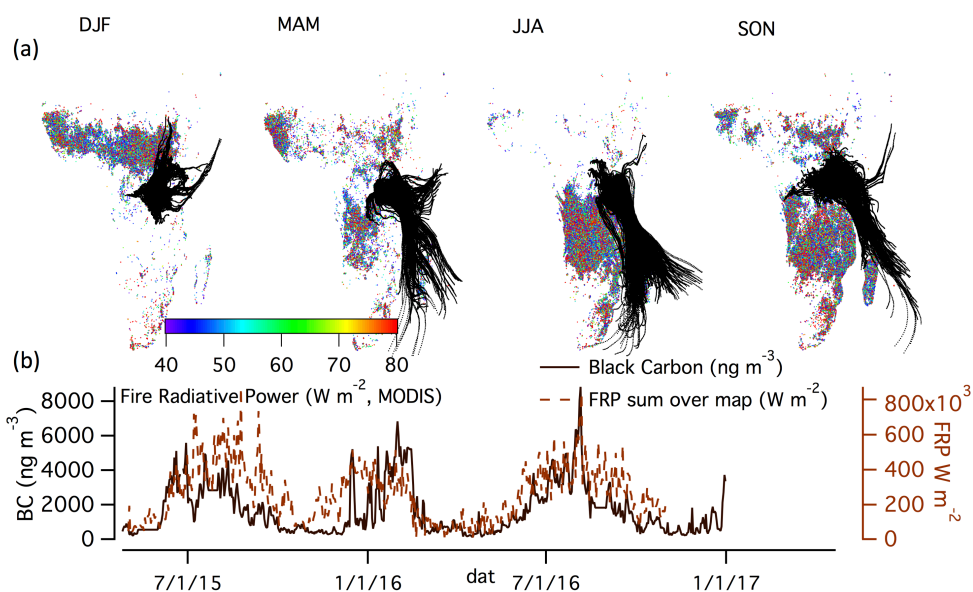
Figure 3. Normalized temporal variations of O₃ mixing ratios, CO mixing ratios, and BC concentrations: (a) diurnal (b) monthly concentrations, and (c) differences by day of the week. Shaded areas are 95% confidence intervals.



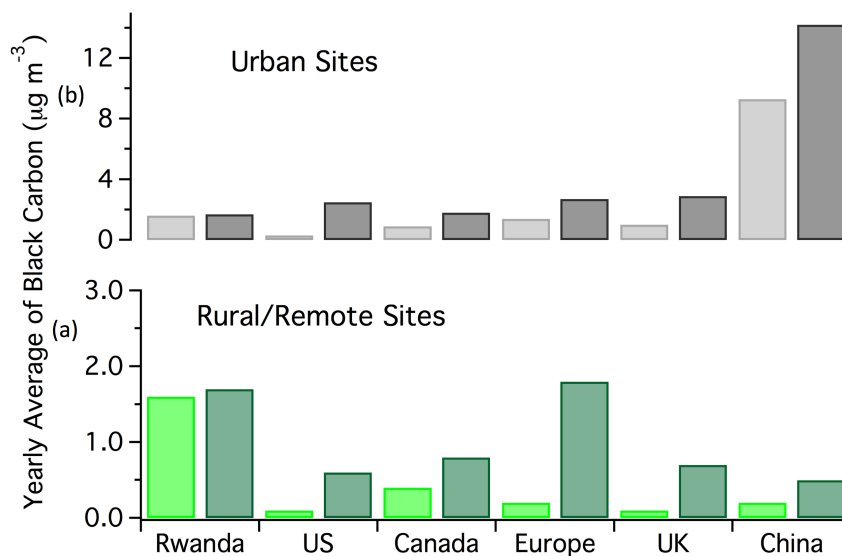
994
995
996
997
998
999

Figure 4: Monthly means of PM_{2.5} concentrations measured at the US Embassies in
1000 (a) Kampala, Uganda and (b) Addis Ababa, Ethiopia (right) from January-December
1001 2016/2017 (as available). Shaded areas are 95% confidence intervals.

1002
1003
1004
1005
1006
1007
1008
1009
1010
1011



1012
1013 **Figure 5.** (a) Seasonal fire radiative power data acquired with the MODIS instrument
1014 and back trajectories of air masses (generated with the HYSPLIT model) reaching the
1015 Rwanda Climate Observatory for the period May 2015 to January 2017. Seasons in
1016 Rwanda are split into: short dry season, December-January-February (DJF), long
1017 rainy season, March-April-May (MAM), long dry season, June-July-August (JJA,) and
1018 short rainy season, September-October-November (SON). (b) The time series of daily
1019 average BC concentration and the daily sum of Fire Radiative Power (W m^{-2}) from the
1020 pictured data bound by the furthest HYSPLIT backtrajectory reaches each season (box
1021 defined by the most north, south, east, and west point the HYSPLIT backtrajectories
1022 reach).
1023



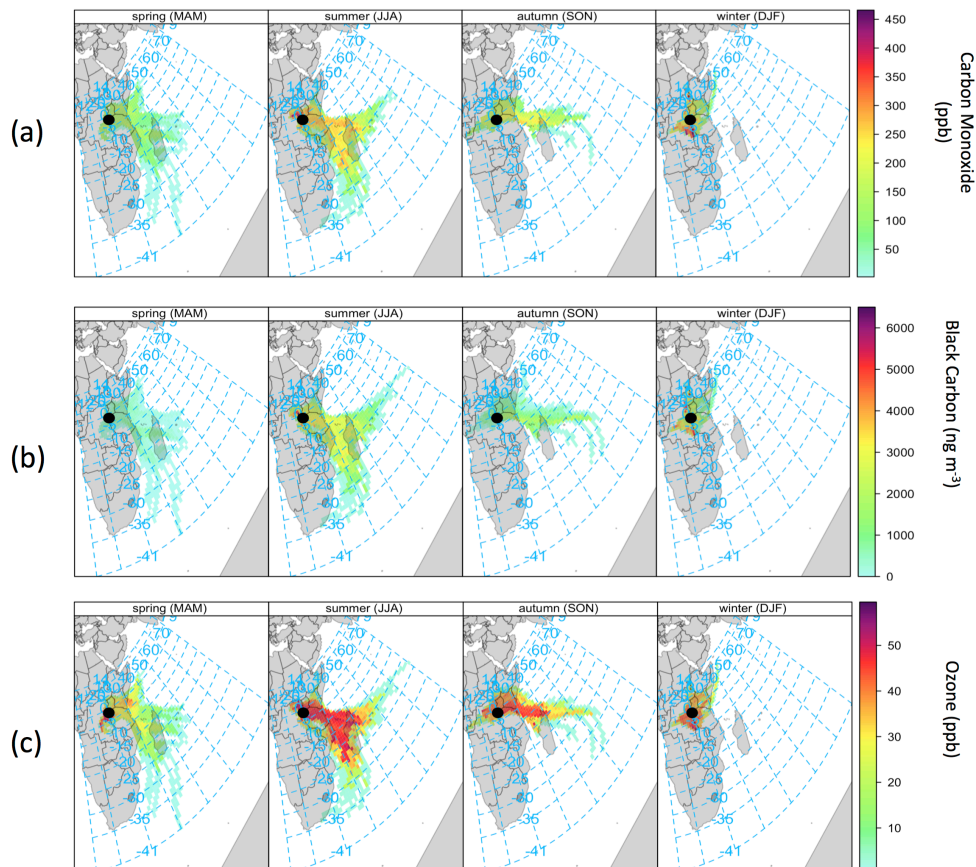
1024
1025 **Figure 6.** (a) Urban and (b) rural maximum (dark grey/green) of annual averages
1026 (dark grey/green) and minimum (light grey/green) of annual averages (light
1027 grey/green) BC concentrations at various sites globally. The BC data for Rwanda is
1028 from one location (Mt. Mugogo, rural), while the data for other locations were from
1029 multiple locations, averaged over one year. The annual average BC concentrations for
1030 Rwanda were calculated for the data from April 1st to April 1st of the next year. There
1031 was BC data for two years measured at RCO. BC data source for other sites:
1032 <https://www3.epa.gov/blackcarbon/2012report/Chapter5.pdf>, compiled from
1033 multiple sources.

1034

1035
1036
1037
1038



1039
1040
1041



1042

1043 **Figure 7.** From top to bottom, concentration-weighted back trajectories of (a) CO, (b)

1044 BC, and (c) O₃, separated by season, for measurements at the Rwanda Climate

1045 Observatory (black dot) for the period of July 2015-January 2017.

1046

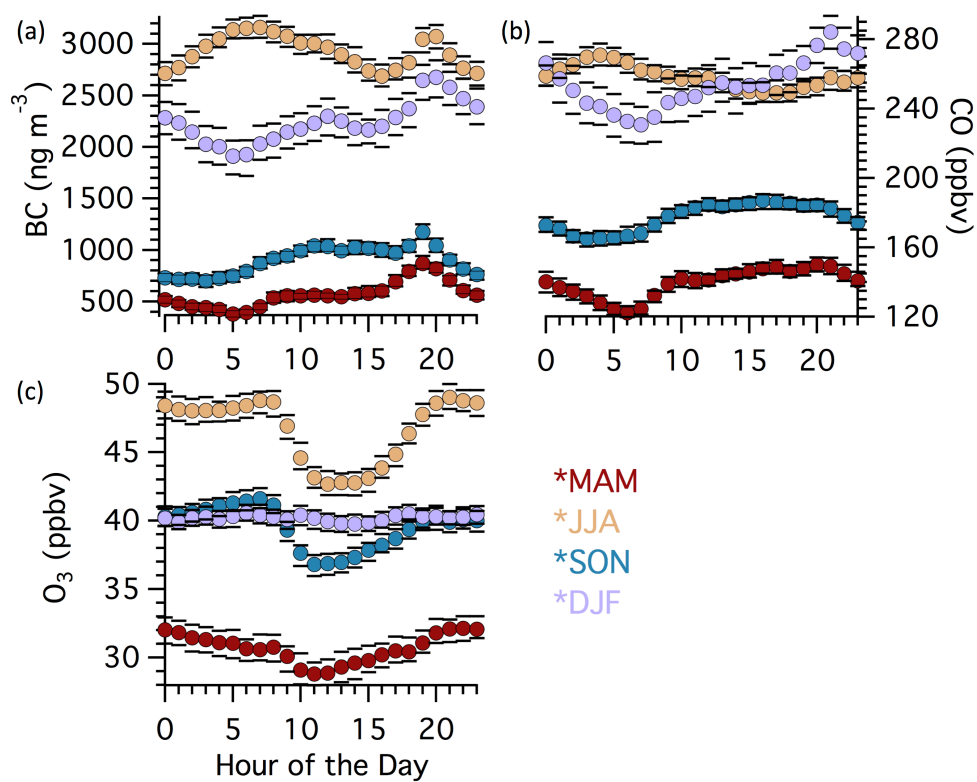
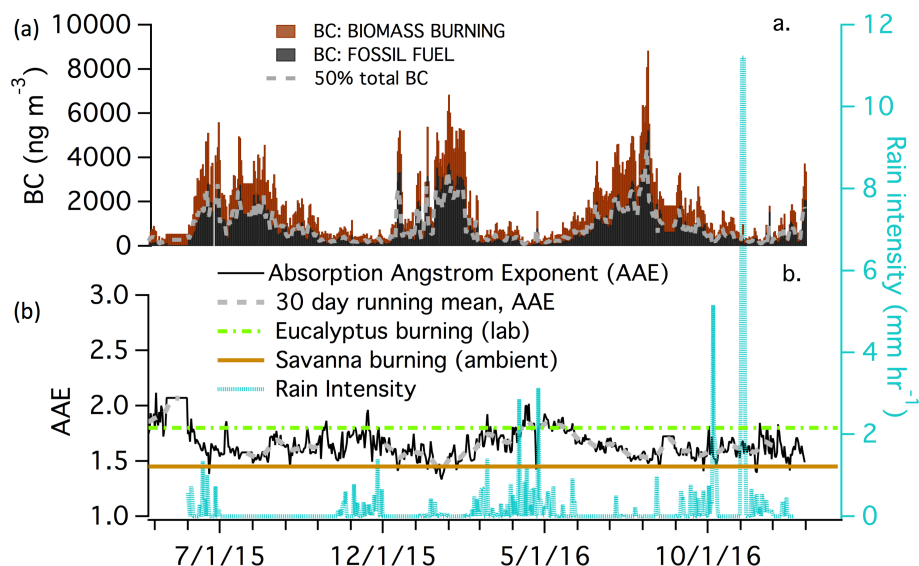


Figure 8. Seasonally separated diurnal profiles of (a) BC concentrations, (b) CO mixing ratios, and (c) O₃ mixing ratios, colored for each season. The circles represent mean concentrations and the lines represent 95% confidence intervals.



1048



1049 **Figure 9.** (a) Time series of contributions of fossil fuel combustion and biomass
1050 burning to BC concentrations observed at RCO. (b) Daily average absorption
1051 Angstrom exponent (AAE) measured at RCO (black line), rain intensity, and published
1052 AAE for Eucalyptus burning ((Yuan et al., 2016), laboratory studies, green line) and
1053 savanna burning ((Russell et al., 2010), ambient, brown line) also shown as reference.
1054