

# 1 **NO<sub>2</sub>-initiated multiphase oxidation of SO<sub>2</sub> by O<sub>2</sub> on CaCO<sub>3</sub> particles**

2 Ting Yu\*, Defeng Zhao\*, Xiaojuan Song, Tong Zhu

3 BIC-ESAT and SKL-ESPC, College of Environmental Sciences and Engineering, Peking University, Beijing,  
4 100871, China

5 \*These authors contributed equally to this work.

6 *Correspondence to:* Tong Zhu (tzhu@pku.edu.cn)

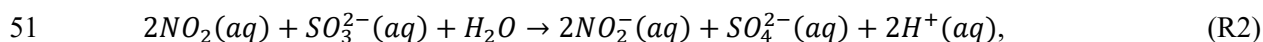
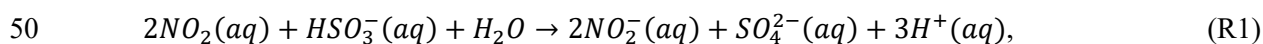
7 **Abstract.** The reaction of SO<sub>2</sub> with NO<sub>2</sub> on the surface of aerosol particles has been suggested to be important in  
8 sulfate formation during severe air pollution episodes in China. However, we found that the direct oxidation of  
9 SO<sub>2</sub> by NO<sub>2</sub> was slow and might not be the main reason for sulfate formation in ambient air. In this study, we  
10 investigated the multiphase reaction of SO<sub>2</sub> with an O<sub>2</sub>/NO<sub>2</sub> mixture on single CaCO<sub>3</sub> particles using  
11 Micro-Raman spectroscopy. The reaction converted the CaCO<sub>3</sub> particle to a Ca(NO<sub>3</sub>)<sub>2</sub> droplet, with CaSO<sub>4</sub>•2H<sub>2</sub>O  
12 solid particles embedded in it, which constituted a significant fraction of the droplet volume at the end of the  
13 reaction. The reactive uptake coefficient of SO<sub>2</sub> for sulfate formation was on the order of 10<sup>-5</sup>, which was higher  
14 than that for the multiphase reaction of SO<sub>2</sub> directly with NO<sub>2</sub> by 2–3 orders of magnitude. According to our  
15 observations and the literature, we found that in the multiphase reaction of SO<sub>2</sub> with the O<sub>2</sub>/NO<sub>2</sub> mixture, O<sub>2</sub> was  
16 the main oxidant of SO<sub>2</sub> and was necessary for radical chain propagation. NO<sub>2</sub> acted as the initiator of radical  
17 formation, but not as the main oxidant. The synergy of NO<sub>2</sub> and O<sub>2</sub> resulted in much faster sulfate formation than  
18 the sum of the reaction rates with NO<sub>2</sub> and with O<sub>2</sub> alone. We estimated that the multiphase oxidation of SO<sub>2</sub> by  
19 O<sub>2</sub> initiated by NO<sub>2</sub> could be an important source of sulfate and a sink of SO<sub>2</sub>, based on the calculated lifetime of  
20 SO<sub>2</sub> regarding the loss through the multiphase reaction versus the loss through the gas-phase reaction with OH  
21 radical. Parameterizing the reactive uptake coefficient of the reaction observed in our laboratory for further model  
22 simulation is needed, as well as an integrated assessment based on field observations, laboratory study results,  
23 and model simulations to evaluate the importance of the reaction in ambient air during severe air pollution  
24 episodes, especially in China.

26 It has been suggested that multiphase or heterogeneous oxidation of SO<sub>2</sub> potentially plays an important role  
 27 in sulfate formation in the atmosphere (Seinfeld and Pandis, 2006). During the severe pollution episodes that  
 28 occur frequently in China, high sulfate concentrations cannot be explained by the gas phase oxidation of SO<sub>2</sub> and  
 29 its well-known aqueous chemistry (Zheng et al., 2015a; Cheng et al., 2016), highlighting the role of  
 30 under-appreciated heterogeneous oxidation or multiphase pathways.

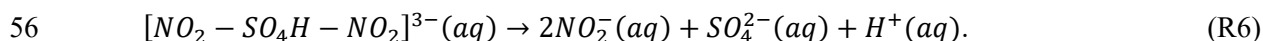
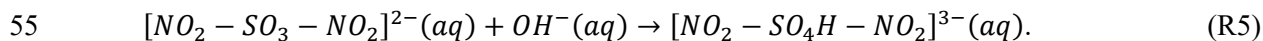
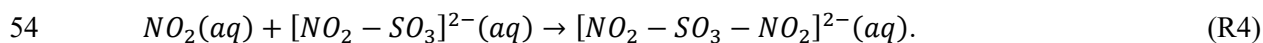
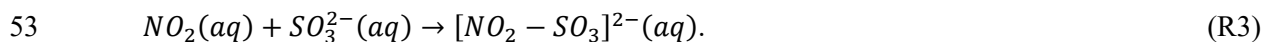
31 Recently, the multiphase oxidation of SO<sub>2</sub> by NO<sub>2</sub> has been introduced in air quality model simulations to  
 32 explain the discrepancy between the modeled and observed sulfate concentration during severe pollution episodes  
 33 in China (Cheng et al., 2016; Gao et al., 2016; Wang et al., 2016; Xue et al., 2016), despite the uncertainties in  
 34 the kinetic parameters for SO<sub>2</sub> oxidation and in the pH value of aerosol particles in China (Wang et al., 2016;  
 35 Cheng et al., 2016; Liu et al., 2017; Guo et al., 2017). However, according to our recently published results (Zhao  
 36 et al., 2017), the direct oxidation of SO<sub>2</sub> by NO<sub>2</sub> could not contribute significantly to sulfate formation in the  
 37 atmosphere because the reactive uptake coefficient of SO<sub>2</sub> for sulfate formation due to direct oxidation by NO<sub>2</sub> is  
 38 very low ( $\sim 10^{-8}$ ).

39 Although the contribution of the direct oxidation of SO<sub>2</sub> by NO<sub>2</sub> to sulfate formation is not significant, NO<sub>2</sub>  
 40 may be involved in other oxidation pathways of SO<sub>2</sub>. It has been reported that the reaction of NO<sub>2</sub> with SO<sub>3</sub><sup>2-</sup> and  
 41 HSO<sub>3</sub><sup>-</sup> in the bulk aqueous phase can form the SO<sub>3</sub><sup>•-</sup> radical, which can further react with O<sub>2</sub> and produce a series  
 42 of radicals that oxidize S(IV) species (Littlejohn et al., 1993). The reaction pathway may result in a fast SO<sub>2</sub>  
 43 oxidation due to the potential synergy of NO<sub>2</sub> and O<sub>2</sub>.

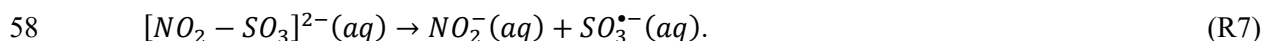
44 Despite such a reaction mechanism for SO<sub>2</sub> oxidation being proposed, its role in SO<sub>2</sub> oxidation in the  
 45 ambient atmosphere is not well established. Most previous studies have focused on the direct reaction of SO<sub>2</sub>  
 46 with NO<sub>2</sub>, including the determination of its rate constant (Lee and Schwartz, 1983; Clifton et al., 1988; Shen and  
 47 Rochelle, 1998; Spindler et al., 2003; Nash, 1979; Huie and Neta, 1986). According to the reaction products and  
 48 their reported yields (Lee and Schwartz, 1983; Clifton et al., 1988), the overall reaction equations of the direct  
 49 reaction of SO<sub>2</sub> with NO<sub>2</sub> are as follows:



52 and the reactions are proposed to proceed via NO<sub>2</sub>-S(IV) adduct complexes (Clifton et al., 1988).



57 Additionally, NO<sub>2</sub>-S(IV) adduct complex may decompose as follows (Spindler et al., 2003).



59 However, studies of the oxidation rate of SO<sub>2</sub> at the O<sub>2</sub> concentrations relevant to the ambient atmosphere and the  
 60 potential influence of the synergy of NO<sub>2</sub> and O<sub>2</sub> on the oxidation rate are very limited (Turšič et al., 2001; He et  
 61 al., 2014), except a few studies investigated SO<sub>2</sub> oxidation in the presence of NO<sub>2</sub> as well as O<sub>2</sub> (Littlejohn et al.,  
 62 1993; Shen and Rochelle, 1998; Santachiara et al., 1990). Moreover, previous studies have mainly focused on the

63 reaction in bulk solution and only few studies have investigated the oxidation of SO<sub>2</sub> by NO<sub>2</sub> on aerosol particles  
64 (Santachiara et al., 1990, 1993). On aerosol particles, water activity, pH, ionic strength, the presence of other  
65 compounds or ions, and the role of particle surface are different from in dilute bulk solution and may affect the  
66 reaction process and reaction rate. Therefore, further studies of the multiphase reaction of SO<sub>2</sub> with O<sub>2</sub>/NO<sub>2</sub>  
67 mixtures on aerosol particles are required to determine the kinetic parameters and the mechanism of the reaction.

68 In this study, we investigated the multiphase reaction of SO<sub>2</sub> with O<sub>2</sub> in the presence of NO<sub>2</sub> on CaCO<sub>3</sub>  
69 particles. We quantified the reactive uptake coefficient of SO<sub>2</sub> due to the reaction with an O<sub>2</sub>/NO<sub>2</sub>/H<sub>2</sub>O mixture.  
70 Based on our observations and the existing literature, we further discussed the reaction mechanism. Furthermore,  
71 we estimated the role of the multiphase oxidation of SO<sub>2</sub> by O<sub>2</sub> in the presence of NO<sub>2</sub> in the atmosphere.

## 72 2 Experimental

73 The experiments were conducted using a flow reaction system and the setup is shown in Fig. S1. The  
74 experimental setup and procedure used have been described in detail in previous studies (Zhao et al., 2017; Zhao  
75 et al., 2011; Liu et al., 2008). A gas mixture of NO<sub>2</sub>, SO<sub>2</sub>, O<sub>2</sub>, N<sub>2</sub>, and water vapor reacted with particles  
76 deposited on a substrate in the flow reaction cell. The concentrations of SO<sub>2</sub> and NO<sub>2</sub> were controlled using mass  
77 flow controllers by varying the flow rates of SO<sub>2</sub> (2,000 ppm in high purity N<sub>2</sub>, National Institute of Metrology  
78 P.R. China), NO<sub>2</sub> (1,000 ppm in high purity N<sub>2</sub>, Messer, Germany), and synthetic air [20% O<sub>2</sub> (high purity grade:  
79 99.999%, Beijing Haikeyuanchang Practical Gas Co., Ltd.) and 80% N<sub>2</sub> (high purity grade: 99.999%, Beijing  
80 Haikeyuanchang Practical Gas Co., Ltd.)]. Relative humidity (RH) was controlled by regulating the flow rates of  
81 reactant gases, dry synthetic air, and humidified synthetic air. Humidified synthetic air was prepared by bubbling  
82 synthetic air through fritted glass in water. In some experiments, the O<sub>2</sub> concentrations were varied by regulating  
83 the mixing ratios of O<sub>2</sub> and N<sub>2</sub> to investigate the effect of O<sub>2</sub>. SO<sub>2</sub>/O<sub>2</sub>/NO<sub>2</sub>/H<sub>2</sub>O mixtures flew through the  
84 reaction cell and reacted with individual stationary CaCO<sub>3</sub> particles, which were deposited on a Teflon-FEP film  
85 substrate annealed to a silicon wafer. RH and temperature were measured using a hygrometer (HMT100, Vaisala,  
86 Vantaa, Finland) at the exit of the reaction cell. Additionally, temperature was measured using another small  
87 temperature sensor (Pt 100, 1/3 DIN B, Heraeus, Hanau, Germany) in the reaction cell. All the experiments were  
88 conducted at 298 ± 0.5 K. The experiments were conducted under two RHs (72% and 82%) at 75 ppm SO<sub>2</sub> and  
89 75 ppm NO<sub>2</sub>.

90 During the reaction, particles were monitored *in-situ* via a glass window on the top of the reaction cell using  
91 a Micro-Raman spectrometer (LabRam HR800, HORIBA Jobin Yvon, Kyoto, Japan) to obtain microscopic  
92 images and Raman spectra. A 514-nm excitation laser was used, and back scattering Raman signals were detected.  
93 The details of the instrument are described elsewhere (Liu et al., 2008; Zhao et al., 2011). Because the particles  
94 were larger than the laser spot in this study (~1.5 μm), confocal Raman mapping was used to measure the spectra  
95 at different locations on a particle to obtain the chemical information of the entire particle. The mapping area was  
96 rectangular and was slightly larger than the particle, with mapping steps of 1 × 1 μm. Raman spectra in the range  
97 of 800–3,900 cm<sup>-1</sup> were acquired with an exposure time of 1 s for each mapping point. Raman spectra were  
98 analyzed using Labspec 5 software (HORIBA Jobin Yvon). Raman peaks were fitted to Gaussian–Lorentzian

99 functions to obtain peak positions and peak areas at different locations on the particle. The peak areas were then  
100 added together to obtain the peak area for the entire particle.

101 Particles of CaCO<sub>3</sub> (98%, Sigma-Aldrich, USA), with average diameters of about 7–10 μm as specified by  
102 the supplier, were used in the experiments. The CaCO<sub>3</sub> particles were rhombohedron crystals; X-ray diffraction  
103 analysis indicated that they were calcite (Fig. S2). Individual particles were prepared by dripping a dilute CaCO<sub>3</sub>  
104 suspended solution onto Teflon-FEP film using a pipette and then drying the sample in an oven at 80°C for 10 h.

105 The amount of CaSO<sub>4</sub> as a reaction product was quantified based on Raman peak areas and particle sizes.  
106 The details of the method are described in our previous study (Zhao et al., 2017). Briefly, the amount of reaction  
107 product CaSO<sub>4</sub> formed was determined as a function of time using Raman peak areas. Raman peak areas were  
108 converted to the amount of compound formed using a calibration curve obtained from pure CaSO<sub>4</sub> particles of  
109 different sizes, which were determined according to microscopic images. The reaction rate, i.e., the sulfate  
110 production rate, was derived from the amount of sulfate formed as a function of time. The reactive uptake  
111 coefficient of SO<sub>2</sub> for sulfate formation ( $\gamma$ ) was further determined from the reaction rate and collision rate of SO<sub>2</sub>  
112 on the surface of a single particle.

$$113 \quad \gamma = \frac{d\{SO_4^{2-}\}}{Z} \cdot dt \quad (1)$$

$$114 \quad Z = \frac{1}{4}cA_s[SO_2], \quad (2)$$

$$115 \quad c = \sqrt{\frac{8RT}{\pi M_{SO_2}}}, \quad (3)$$

116 where R is the gas constant, T is temperature, M<sub>SO<sub>2</sub></sub> is the molecular weight of SO<sub>2</sub>, c is the mean molecular  
117 velocity of SO<sub>2</sub>, A<sub>s</sub> is the surface area of an individual particle, and Z is the collision rate of SO<sub>2</sub> on the surface of  
118 a particle. {SO<sub>4</sub><sup>2-</sup>} indicates the amount of sulfate in the particle phase in moles. The average reaction rate and  
119 surface area of particles during the multiphase reaction period were used to derive the reactive uptake coefficient.  
120 The period was chosen to start after the induction period when ~10% of the final sulfate was formed. [SO<sub>2</sub>]  
121 indicates the concentration of SO<sub>2</sub> in the gas phase.

122 The influence of gas phase diffusion on reactive uptake was evaluated using the resistor model described by  
123 Davidovits et al. (2006) and references therein, as well as using the gas phase diffusion correction factor for a  
124 reactive uptake coefficient according to the method described by Pöschl et al. (2007). The reactive uptake of SO<sub>2</sub>  
125 was found to not be limited by gas phase diffusion (see details in the Supplement S1).

126 In addition, we conducted experiments of the reaction SO<sub>2</sub> with only O<sub>2</sub> on both CaCO<sub>3</sub> solid particles and  
127 internally mixed CaCO<sub>3</sub>/Ca(NO<sub>3</sub>)<sub>2</sub> particles (with CaCO<sub>3</sub> embedded in Ca(NO<sub>3</sub>)<sub>2</sub> droplets), while keeping other  
128 conditions the same as the reaction of SO<sub>2</sub> with an O<sub>2</sub>/NO<sub>2</sub> mixture. These experiments of the multiphase  
129 oxidation of SO<sub>2</sub> by O<sub>2</sub> can help determine the role of NO<sub>2</sub> in the reaction of SO<sub>2</sub> with an O<sub>2</sub>/NO<sub>2</sub> mixture.

### 130 3 Results and discussion

#### 131 3.1 Reaction products and changes in particle morphology

132 Figure 1 shows the Raman spectra of a CaCO<sub>3</sub> particle during the multiphase reaction of SO<sub>2</sub> with O<sub>2</sub>/NO<sub>2</sub>/H<sub>2</sub>O  
133 on its surface. The peak at 1,087 cm<sup>-1</sup> was assigned to the symmetric stretching of carbonate ( $\nu_s(\text{CO}_3^{2-})$ )

134 (Nakamoto, 1997). During the reaction, the peak at  $1,087\text{ cm}^{-1}$  decreased continuously and finally disappeared as  
135 new peaks were observed. The peak at  $1,050\text{ cm}^{-1}$  was assigned to the symmetric stretching of nitrate ( $\nu_s(\text{NO}_3^-)$ ).  
136 The peaks at  $1,010\text{ cm}^{-1}$  and  $1,136\text{ cm}^{-1}$  were assigned to the symmetric stretching ( $\nu_s(\text{SO}_4^{2-})$ ) and asymmetric  
137 stretching ( $\nu_{as}(\text{SO}_4^{2-})$ ) of sulfate in gypsum ( $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$ ), respectively (Sarma et al., 1998). In addition, after the  
138 reaction, a broad envelope in the range of  $2,800\text{--}3,800\text{ cm}^{-1}$  assigned to the stretching of the OH bond in water  
139 molecules was observed. Above this envelope, there were two peaks at  $3,408\text{ cm}^{-1}$  and  $3,497\text{ cm}^{-1}$ , which were  
140 assigned to OH bond stretching in crystallization water of  $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$  (Sarma et al., 1998; Ma et al., 2013).

141 During the multiphase reaction with the  $\text{SO}_2/\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$  mixture, the  $\text{CaCO}_3$  particles displayed a remarkable  
142 change in morphology. The original  $\text{CaCO}_3$  particle was a rhombohedron crystal (Fig. 2, panel i, a). As the  
143 reaction proceeded, its edges became smoother and later a transparent droplet layer formed, which had a newly  
144 formed solid phase embed in it (Fig. 2, panel i, d). The size of the new solid phase grew during the reaction (Fig.  
145 2, panel i, d–f) and it seemed to contain many micro-crystals. Raman mapping revealed that the new solid phase  
146 consisted of  $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$  (Fig. 2, panel iv), and the surrounding aqueous layer consisted of  $\text{Ca}(\text{NO}_3)_2$  (Fig. 2,  
147 panel iii).

148 The particle morphology change shown in Fig. 2 was significantly different from the morphology change in  
149 the direct reaction of  $\text{SO}_2$  with  $\text{NO}_2$  (Zhao et al., 2017), where the  $\text{CaCO}_3$  particle was first converted to a  
150 spherical  $\text{Ca}(\text{NO}_3)_2$  droplet and then needle-shaped  $\text{CaSO}_4$  crystals formed inside the droplet (Zhao et al., 2017).  
151 Moreover, the amount of  $\text{CaSO}_4$  formed in this study was much higher than that in the direct reaction of  $\text{SO}_2$  with  
152  $\text{NO}_2$ . The  $\text{CaSO}_4$  solid particle constituted a significant fraction of the volume of the droplet, while in the direct  
153 reaction of  $\text{SO}_2$  with  $\text{NO}_2$  the few needle-shaped  $\text{CaSO}_4$  crystals that formed only constituted a small fraction of  
154 the droplet volume (Zhao et al., 2017).

### 155 3.2 Reaction process

156 During the reaction, the amounts of carbonate, nitrate, and sulfate were determined as a function of time, as  
157 shown in Fig. 3. At the beginning of the reaction, the amount of carbonate decreased slowly, while the amount of  
158 nitrate and sulfate increased slowly. After a period of induction of around 50 min, the reaction accelerated  
159 significantly, leading to a rapid consumption of carbonate and production of nitrate and sulfate. The decrease in  
160 the amount of carbonate and the increase in the amount of nitrate was because carbonate reacted continuously  
161 with  $\text{NO}_2$  and  $\text{H}_2\text{O}$ , forming  $\text{Ca}(\text{NO}_3)_2$ . The detailed mechanism of the multiphase reaction of carbonate with  $\text{NO}_2$   
162 and  $\text{H}_2\text{O}$  were discussed in our previous studies (Li et al., 2010; Zhao et al., 2017). The mechanism of sulfate  
163 formation is discussed in detail in Section 3.4 of the present study. Finally, the carbonate was completely  
164 consumed, and the amounts of nitrate and sulfate levelled off.

165 Figure 3 shows that nitrate and sulfate were formed simultaneously during the reaction. This contrasts with  
166 the observations made during the direct reaction of  $\text{SO}_2$  with  $\text{NO}_2$ , where nitrate was formed first, and sulfate was  
167 essentially formed after the complete conversion of  $\text{CaCO}_3$  particles to  $\text{Ca}(\text{NO}_3)_2$  droplets (Zhao et al., 2017).  
168 Moreover, the time taken for carbonate to be completely consumed was longer in this study than in the direct  
169 reaction of  $\text{SO}_2$  with  $\text{NO}_2$  ( $\sim 120$  vs.  $\sim 40$  min) when other conditions were kept the same (Zhao et al., 2017).

### 170 3.3 Reactive uptake coefficient of SO<sub>2</sub>

171 The reactive uptake coefficients of SO<sub>2</sub> for sulfate formation ( $\gamma$ ) in the reaction of SO<sub>2</sub> with the  
172 O<sub>2</sub>/NO<sub>2</sub>/H<sub>2</sub>O/N<sub>2</sub> mixture on CaCO<sub>3</sub> with various O<sub>2</sub> concentrations are shown in Table 1. The value of  $\gamma$  for the  
173 reaction of SO<sub>2</sub> with O<sub>2</sub>/NO<sub>2</sub> at three O<sub>2</sub> concentrations (5, 20, and 86%) was in the range of  $(0.35\text{--}1.7) \times 10^{-5}$ ,  
174 and was  $1.2 \times 10^{-5}$  in synthetic air. This latter value was 2–3 orders of magnitude higher than that for the reaction  
175 of SO<sub>2</sub> directly with NO<sub>2</sub> under similar conditions (Zhao et al., 2017). When other conditions were kept constant,  
176  $\gamma$  increased with the O<sub>2</sub> concentration. This indicates that O<sub>2</sub> played a key role in enhancing the oxidation rate of  
177 SO<sub>2</sub>.

178 The role of O<sub>2</sub> in enhancing the reactive uptake of SO<sub>2</sub> reported here is consistent with the findings in some  
179 previous studies. For example, Littlejohn et al. (1993)'s data showed that sulfite oxidation rate increases with the  
180 O<sub>2</sub> concentration (0–5% by volume). Shen and Rochelle (1998) also found that in the presence of O<sub>2</sub>, the aqueous  
181 sulfite oxidation rate is enhanced. By investigating the oxidation of SO<sub>2</sub> by NO<sub>2</sub> in monodispersed water droplets  
182 growing on carbon nuclei, Santachiara et al. (1990) found that sulfate formation rate with 2% O<sub>2</sub> is much higher  
183 than that without O<sub>2</sub>. Yet, our findings, as well as those in the studies referred to above, are in contrast to those  
184 reported by Lee and Schwartz (1983), who found that changing from N<sub>2</sub> to air as a carrier gas only increases  
185 bisulfite oxidation rate by no more than 10%. The difference between our study and Lee and Schwartz (1983)  
186 could be due to the difference in O<sub>2</sub> diffusion from gas to the condensed phase and the different mechanisms  
187 between the multiphase reaction on particles and the aqueous reaction.

188 Only few studies have reported the S(IV) oxidation rate in the reaction of S(IV) with O<sub>2</sub>/NO<sub>2</sub> mixtures  
189 (Turšič et al., 2001; Littlejohn et al., 1993). However, due to the limiting step by the aqueous phase mass transfer,  
190 it is difficult to quantitatively compare the reaction rates in those studies with the uptake coefficient in our study  
191 and the rate constants determined by Lee and Schwartz (1983) and Clifton et al. (1988). For example, a rate  
192 constant of  $2.4 \times 10^3 \text{ mol}^{-1} \text{ L s}^{-1}$  (at pH 3) can be derived from the results of Turšič et al. (2001), which is much  
193 lower than the values reported by Lee and Schwartz (1983) and Clifton et al. (1988). This can be attributed to the  
194 limiting step by the aqueous-phase mass transfer because the characteristic mixing time in the aqueous phase in  
195 Turšič et al. (2001) was likely much longer than that of Lee and Schwartz (1983) (1.7–5.3 s), according to the  
196 HSO<sub>3</sub><sup>-</sup> concentration time series reported by Turšič et al. (2001).

197 It is important to note that the concentrations of NO<sub>2</sub> and SO<sub>2</sub> used in this study are much higher than those  
198 in the ambient atmosphere. High concentrations of reactant gases are often used in laboratory studies in order to  
199 simulate the ambient reactions at the time scale of days or weeks and to get high signal-to-noise ratios for  
200 detecting products within minutes or hours. In the ambient atmosphere, reactive uptake coefficient of SO<sub>2</sub> should  
201 be lower than that in this study due to the lower NO<sub>2</sub> concentrations when other conditions are comparable and  
202 the chemical/physical processes observed in this study, such as changes in particle composition, phase,  
203 hygroscopicity, and pH should be much slower due to the lower concentrations of NO<sub>2</sub> and SO<sub>2</sub>.

### 204 3.4 Reaction mechanism

205 In the multiphase reaction of SO<sub>2</sub> with O<sub>2</sub>/NO<sub>2</sub>/H<sub>2</sub>O on CaCO<sub>3</sub> particles, we found that CaCO<sub>3</sub> reacted with  
206 NO<sub>2</sub> and H<sub>2</sub>O and produced Ca(NO<sub>3</sub>)<sub>2</sub>, which deliquesced, forming liquid water, and provided a site for the

207 aqueous oxidation of SO<sub>2</sub>. This process is similar to the direct reaction of SO<sub>2</sub> with NO<sub>2</sub> on CaCO<sub>3</sub> particles. The  
208 details of this part of the reaction mechanism were discussed in our previous study (Zhao et al., 2017).

209 Once the aqueous phase was formed, SO<sub>2</sub> could undergo multiphase reactions with O<sub>2</sub>/NO<sub>2</sub>. The mechanism  
210 of the direct aqueous reaction of S(IV) with NO<sub>2</sub> in the absence of O<sub>2</sub> is complex. Previous studies have proposed  
211 two different mechanisms for the reaction. One involves SO<sub>3</sub><sup>•-</sup> radical formation (Littlejohn et al., 1993; Shen and  
212 Rochelle, 1998; Turšič et al., 2001) (referred as “free-radical chain” mechanism, while the other involves the  
213 formation of NO<sub>2</sub>-S(IV) complexes (Clifton et al., 1988), but no radical formation (referred as “NO<sub>2</sub>-S(IV)  
214 complex” mechanism).

215 According to the NO<sub>2</sub>-S(IV) complex mechanism, the presence of O<sub>2</sub> should not affect the SO<sub>2</sub> oxidation  
216 rate; however, in this study, a substantial enhancement in the SO<sub>2</sub> oxidation rate was observed in the presence of  
217 O<sub>2</sub> compared with that in the absence of O<sub>2</sub>. Therefore, the NO<sub>2</sub>-S(IV) complex mechanism was less likely to  
218 have been important in this study.

219 In the free-radical chain mechanism, the SO<sub>3</sub><sup>•-</sup> radical is proposed to be formed (R8, Table 2), which is based  
220 on the observation of S<sub>2</sub>O<sub>6</sub><sup>2-</sup> formation, with S<sub>2</sub>O<sub>6</sub><sup>2-</sup> known to be the combination reaction product of SO<sub>3</sub><sup>•-</sup>  
221 (Eriksen, 1974; Hayon et al., 1972; Deister and Warneck, 1990; Brandt et al., 1994; Waygood and McElroy,  
222 1992). In addition to SO<sub>4</sub><sup>2-</sup> and NO<sub>2</sub><sup>-</sup>, S<sub>2</sub>O<sub>6</sub><sup>2-</sup> was detected with an appreciable yield using Raman spectroscopy,  
223 following the reaction of NO<sub>2</sub> with aqueous sulfite (Littlejohn et al., 1993). S<sub>2</sub>O<sub>6</sub><sup>2-</sup> was also observed in the  
224 aqueous oxidation of bisulfite in an N<sub>2</sub>-saturated solution in the presence of Fe(III) using ion-interaction  
225 chromatography (Podkrajšek et al., 2002). The SO<sub>3</sub><sup>•-</sup> radical can react via two pathways, forming either S<sub>2</sub>O<sub>6</sub><sup>2-</sup> or  
226 SO<sub>4</sub><sup>2-</sup> (R9–R11, Table 2). The reactions R9–R11 have been well established in studies of S(IV) oxidation by  
227 other pathways, including OH oxidation, photo-oxidation, and transition metal catalyzed oxidation (Eriksen, 1974;  
228 Hayon et al., 1972; Deister and Warneck, 1990; Brandt et al., 1994; Brandt and Vaneldik, 1995; Waygood and  
229 McElroy, 1992). In addition, although previous studies have not reported the direct observation of the SO<sub>3</sub><sup>•-</sup>  
230 radical in the aqueous reaction of S(IV) with NO<sub>2</sub>, SO<sub>3</sub><sup>•-</sup> was observed in the reaction of NO<sub>2</sub><sup>-</sup> with SO<sub>3</sub><sup>2-</sup> in an  
231 acidic buffer solution (pH = 4.0) using electron spin resonance (ESR) (Shi, 1994). Because NO<sub>2</sub><sup>-</sup> is formed in the  
232 aqueous reaction of SO<sub>2</sub> with NO<sub>2</sub>, and S<sub>2</sub>O<sub>6</sub><sup>2-</sup> as the combination reaction product of SO<sub>3</sub><sup>•-</sup> is observed (Littlejohn  
233 et al., 1993), SO<sub>3</sub><sup>•-</sup> formation is plausible.

234 In the presence of O<sub>2</sub>, the SO<sub>3</sub><sup>•-</sup> radical can react rapidly with O<sub>2</sub>, forming the SO<sub>5</sub><sup>•-</sup> radical (R12, Table 2).  
235 Following this reaction, a number of chain reactions can occur to ultimately form sulfate (Littlejohn et al., 1993;  
236 Seinfeld and Pandis, 2006; Shen and Rochelle, 1998) (R13–R16, Table 2). Littlejohn et al. (1993) observed that  
237 the amount of S<sub>2</sub>O<sub>6</sub><sup>2-</sup> relative to SO<sub>4</sub><sup>2-</sup> formed in the aqueous reaction of NO<sub>2</sub> with sulfite decreases in the  
238 presence of O<sub>2</sub> compared with the reaction in the absence of O<sub>2</sub>. At low NO<sub>2</sub> concentrations (< 5 ppm), S<sub>2</sub>O<sub>6</sub><sup>2-</sup> is  
239 undetectable in the presence of O<sub>2</sub>. This indicates that O<sub>2</sub> suppresses the reaction pathway of S<sub>2</sub>O<sub>6</sub><sup>2-</sup> formation  
240 (R9, Table 2). Because the SO<sub>3</sub><sup>•-</sup> radical can react rapidly with O<sub>2</sub>, forming the SO<sub>5</sub><sup>•-</sup> radical, and would therefore  
241 be consumed, the suppression of S<sub>2</sub>O<sub>6</sub><sup>2-</sup> formation can be attributed to the reaction of SO<sub>3</sub><sup>•-</sup> with O<sub>2</sub> (R12, Table  
242 2). The reactions R12–R16 have been well established by studies of the oxidation of S(IV) by OH or  
243 photo-oxidation, and all the radicals have been observed (Hayon et al., 1972; Huie et al., 1989; Huie and Neta,  
244 1987; Chameides and Davis, 1982; Seinfeld and Pandis, 2006).

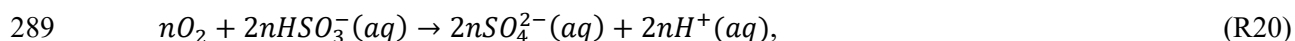
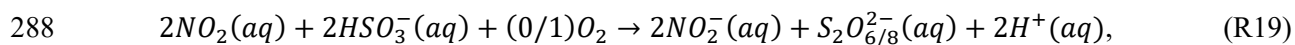
245 The free-radical chain mechanism is consistent with the findings of this study and is therefore more plausible.  
246 The enhancement of the SO<sub>2</sub> oxidation rate in the reaction of SO<sub>2</sub> with O<sub>2</sub>/NO<sub>2</sub>/H<sub>2</sub>O on CaCO<sub>3</sub> particles  
247 compared with that in the direct reaction of SO<sub>2</sub> with NO<sub>2</sub>/H<sub>2</sub>O was attributed to O<sub>2</sub>. Although during the reaction  
248 in the absence of O<sub>2</sub>—i.e., the direct oxidation of SO<sub>2</sub> by NO<sub>2</sub>—the SO<sub>3</sub><sup>•-</sup> radical can be formed (R7), the reaction  
249 chain cannot propagate (R12–R16). Therefore, the S(IV) oxidation rate and the reactive uptake coefficient of SO<sub>2</sub>  
250 were much lower than that in the presence of O<sub>2</sub>. According to the difference between the reactive uptake  
251 coefficient in this study and in the direct reaction of SO<sub>2</sub> with NO<sub>2</sub> (Zhao et al., 2017), the sulfate production rate  
252 via chain reactions due to the presence of O<sub>2</sub> (20%) was 2–3 orders of magnitude faster than the direct oxidation  
253 of SO<sub>2</sub> by NO<sub>2</sub>. This indicates that sulfate production in the reaction of SO<sub>2</sub> with O<sub>2</sub>/NO<sub>2</sub> was largely due to O<sub>2</sub>  
254 oxidation via the chain reaction pathway, i.e., “autoxidation” of S(IV), rather than the direct oxidation of SO<sub>2</sub> by  
255 NO<sub>2</sub> and thus O<sub>2</sub> was the main oxidant of SO<sub>2</sub>.

256 In addition to the two mechanisms above, Spindler et al. (2003) proposed a reaction mechanism involving  
257 first NO<sub>2</sub>–S(IV) complex formation and subsequent SO<sub>3</sub><sup>•-</sup> radical formation (R3, R7). NO<sub>2</sub>–S(IV) complex may  
258 establish an equilibrium with SO<sub>3</sub><sup>•-</sup> in contrast to the direct formation of SO<sub>3</sub><sup>•-</sup> via the reaction of NO<sub>2</sub> with SO<sub>2</sub>.  
259 Higher concentration of O<sub>2</sub> favors the conversion of SO<sub>3</sub><sup>•-</sup> to SO<sub>5</sub><sup>•-</sup> and thus can reduce the SO<sub>3</sub><sup>•-</sup> concentration,  
260 shifting the equilibrium to the product side and promoting the overall S(IV) oxidation. O<sub>2</sub> can act in a similar way  
261 as in the free-radical chain mechanism. Admittedly, we cannot rule out the possibility NO<sub>2</sub>–S(IV) complex  
262 formation. But such a mechanism can still be classified as the free-radical chain mechanism since the S(IV)  
263 oxidation still proceeds via the radical chain reactions. Although the direct oxidation of SO<sub>2</sub> by NO<sub>2</sub> only  
264 accounted for a very small fraction of sulfate formation, NO<sub>2</sub> played an important role in the SO<sub>2</sub> oxidation by  
265 initiating the chain reactions via the production of the SO<sub>3</sub><sup>•-</sup> radical (R7). In the experiment without NO<sub>2</sub>, but with  
266 other reaction conditions the same, we were unable to detect sulfate after 5 h of reaction. This indicates that O<sub>2</sub> by  
267 itself cannot initiate the chain reactions (although it favors chain propagation), and that the oxidation of SO<sub>2</sub> by  
268 O<sub>2</sub> was slow. The effect on the SO<sub>2</sub> oxidation rate when both NO<sub>2</sub> and O<sub>2</sub> were present was much higher than the  
269 sum of the effect of NO<sub>2</sub> and O<sub>2</sub>. We refer to this effect as the synergy of NO<sub>2</sub> and O<sub>2</sub>, which resulted in the fast  
270 oxidation of SO<sub>2</sub> to form sulfate in this study. This effect is similar to a “ternary” reaction found with the reaction  
271 of NO<sub>2</sub>–particles–H<sub>2</sub>O or SO<sub>2</sub>–particles–O<sub>3</sub> (Zhu et al., 2011), where the reaction rate can be much faster than the  
272 sum of the reaction rates for the reaction of the second and third reactant with the first reactant. In addition to  
273 acting as the initiator of chain reactions, NO<sub>2</sub> also contributed to the formation of the aqueous phase through the  
274 reaction with CaCO<sub>3</sub>, forming Ca(NO<sub>3</sub>)<sub>2</sub> as discussed above, which provided a site for S(IV) oxidation.

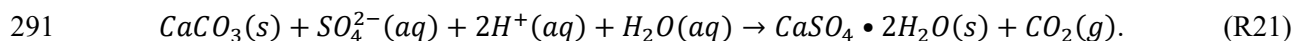
275 Based on the discussion above, we summarize the reaction mechanism that occurred in this study in Table 2.  
276 The reactions are classified as chain initiation, chain propagation, and chain termination. The dominant S(IV)  
277 species depends on pH. Due to the fast dissociations of SO<sub>2</sub>•H<sub>2</sub>O and HSO<sub>3</sub><sup>-</sup>, reactions consuming one of these  
278 S(IV) species will result in instantaneous re-establishment of the equilibria between them (Seinfeld and Pandis,  
279 2006). In this study, the pH of the aqueous layer of Ca(NO<sub>3</sub>)<sub>2</sub> may change dynamically with time during the  
280 reaction and may not be completely homogeneous within the aqueous droplet. The pH values could vary between  
281 ~3 and ~7.6. In the surface of the aqueous layer, pH was mainly determined by the gas–aqueous equilibrium of  
282 SO<sub>2</sub>, and was estimated to be ~3. In the vicinity of the CaCO<sub>3</sub> core, pH was mainly determined by the hydrolysis  
283 of carbonate and was estimated to be ~7.6. It is likely that both HSO<sub>3</sub><sup>-</sup> and SO<sub>3</sub><sup>2-</sup> were present, and the dominant



284 species depended on the reaction time and location within the aqueous droplet. Nevertheless, to make the reaction  
285 mechanism clearer,  $\text{HSO}_3^-$  was used in the reaction equations. Similar reaction equations are also applicable to  
286  $\text{SO}_3^{2-}$  because of the fast dissociations of  $\text{SO}_2 \cdot \text{H}_2\text{O}$  and  $\text{HSO}_3^-$ . Overall, the reaction can be written as follows,  
287 which clearly shows that  $\text{O}_2$  was the main oxidant for sulfate formation:



290 where  $n \gg 1$ . Once sulfuric acid was formed, it reacted with  $\text{CaCO}_3$ , forming  $\text{CaSO}_4$ :



292 As mentioned above, compared with the direct reaction of  $\text{SO}_2$  with  $\text{NO}_2$ ,  $\text{CaCO}_3$  was consumed more  
293 slowly in the reaction with  $\text{O}_2/\text{NO}_2$ . There were two possible reasons for this. First, the  $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$  formed in  
294 the reaction could cover the  $\text{CaCO}_3$  surface and partly suppress the diffusion of aqueous ions, such as protons,  
295 and also limit the contact of reactants with the surface of the  $\text{CaCO}_3$  particles, thus reducing the  $\text{CaCO}_3$   
296 consumption rate. Second, compared with the direct reaction of  $\text{SO}_2$  with  $\text{NO}_2$ , a much higher fraction of  $\text{CaCO}_3$   
297 was converted to  $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$  instead of  $\text{Ca}(\text{NO}_3)_2$  due to the fast production of  $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$ . Therefore, the  
298 volume of a  $\text{Ca}(\text{NO}_3)_2$  droplet was much smaller than that in the direct reaction of  $\text{SO}_2$  with  $\text{NO}_2$  for a given  
299  $\text{CaCO}_3$  particle. Because the uptake rate of  $\text{NO}_2$  was proportional to the droplet surface area and the  $\text{NO}_2$   
300 hydrolysis rate was proportional to the droplet volume, the rate of nitric acid production from  $\text{NO}_2$  hydrolysis and  
301 its reaction rate with  $\text{CaCO}_3$  were reduced. Therefore, the  $\text{CaCO}_3$  particles were consumed more slowly in the  
302 reaction with  $\text{O}_2/\text{NO}_2$ .

#### 303 4 Conclusion and implications

304 We investigated the multiphase reaction of  $\text{SO}_2$  with  $\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$  on  $\text{CaCO}_3$  particles. The reaction  
305 converted  $\text{CaCO}_3$  particles to  $\text{Ca}(\text{NO}_3)_2$  droplets, in which  $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$  was embedded and accounted for a  
306 significant fraction of the droplet volume by the end of the reaction. The  $\text{Ca}(\text{NO}_3)_2$  droplet formed by the reaction  
307 of  $\text{CaCO}_3$  with  $\text{NO}_2$  provided a site for the multiphase oxidation of  $\text{SO}_2$ . Generally, nitrate and sulfate were  
308 formed simultaneously. The reactive uptake coefficient of  $\text{SO}_2$  for sulfate formation in the reaction of  $\text{SO}_2$  with  
309  $\text{NO}_2/\text{H}_2\text{O}$  in synthetic air was determined to be around  $10^{-5}$ . Compared with the reaction of  $\text{SO}_2$  with  $\text{NO}_2$  on a  
310  $\text{CaCO}_3$  particle in the absence of  $\text{O}_2$ , i.e., the direct oxidation of  $\text{SO}_2$  by  $\text{NO}_2$  in  $\text{N}_2$ , sulfate production rate in the  
311 reaction of  $\text{SO}_2$  with  $\text{O}_2/\text{NO}_2$  was enhanced by 2–3 orders of magnitude. According to the findings of this study  
312 and the existing literature,  $\text{SO}_2$  oxidation likely proceeded via a free-radical chain reaction mechanism.  $\text{O}_2$  was  
313 the main oxidant of  $\text{SO}_2$ , and  $\text{NO}_2$  mainly acted as an initiator of the chain reactions. The synergy of  $\text{NO}_2$  and  $\text{O}_2$   
314 resulted in the fast oxidation of  $\text{SO}_2$ . The absence of either  $\text{NO}_2$  or  $\text{O}_2$  led to much slower  $\text{SO}_2$  oxidation.

315 Using a method developed in our previous study (Zhao et al., 2017), we assessed the importance of the  
316 multiphase oxidation of  $\text{SO}_2$  by  $\text{O}_2$  in the presence of  $\text{NO}_2$  by estimating the lifetime of  $\text{SO}_2$  due to multiphase  
317 reactions and the lifetime due to the gas phase reaction (with the OH radical). The lifetime of  $\text{SO}_2$  due to the  
318 multiphase reaction of  $\text{SO}_2$  with  $\text{O}_2/\text{NO}_2$  was estimated to be around 20 days using the reactive uptake coefficient  
319 of  $\text{SO}_2$  ( $1.2 \times 10^{-5}$ ) and the typical particle surface area concentration for mineral aerosols in winter in Beijing  
320 ( $6.3 \times 10^{-6} \text{ cm}^2 \text{ cm}^{-3}$ ) (Huang et al., 2015). This lifetime is comparable to the lifetime of  $\text{SO}_2$  due to the gas phase

321 reaction with OH, which is ~12 days assuming that the daytime OH concentration is  $1 \times 10^6$  molecules  $\text{cm}^{-3}$   
322 (Lelieveld et al., 2016; Prinn et al., 2005). Therefore, we conclude that the multiphase oxidation of  $\text{SO}_2$  by  $\text{O}_2$  in  
323 the presence of  $\text{NO}_2$  is likely to be an important source of sulfate and a sink of  $\text{SO}_2$  in the ambient atmosphere,  
324 and can play a significant role in the sulfate formation during severe haze episodes, such as those that frequently  
325 occur in China. During haze episodes, there are high concentrations of  $\text{SO}_2$  and  $\text{NO}_2$  and relative humidity is  
326 often high (Zhang et al., 2014; Wang et al., 2016; Zheng et al., 2015b). Under these conditions, the multiphase  
327 oxidation of  $\text{SO}_2$  by  $\text{O}_2$  in the presence of  $\text{NO}_2$  could proceed rapidly, forming sulfate. The enhanced sulfate  
328 concentration due to multiphase reactions and resulting aerosol water content can further promote the multiphase  
329 oxidation of  $\text{SO}_2$ . The reaction thus proceeds in a self-accelerating way. Therefore, it can contribute significantly  
330 to sulfate formation during haze episodes, which could explain the discrepancies between the observed and  
331 modelled sulfate concentrations (Cheng et al., 2016; Gao et al., 2016; Wang et al., 2016; Zheng et al., 2015a).

332 In addition, elucidating the mechanism of the multiphase reaction of  $\text{SO}_2$  with  $\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$  in the  
333 atmosphere is important for the other atmospheric implications of the reaction besides sulfate formation.  
334 According to the reaction mechanism, the direct oxidation of  $\text{SO}_2$  by  $\text{NO}_2$  forms sulfate and nitrite, with a  
335 stoichiometry of 1:1, and nitrite can further form HONO under acidic conditions. The HONO could then  
336 evaporate into the atmosphere, where it would be an important source of OH radical. If  $\text{NO}_2$  were the main  
337 oxidant of  $\text{SO}_2$  in the multiphase reaction, the reaction would form one HONO molecule for every sulfate formed.  
338 Thus, the oxidation of  $\text{SO}_2$  by  $\text{NO}_2$  can simultaneously be an important source of HONO and OH radical, and  
339  $\text{SO}_2$  oxidation would be strongly coupled with reactive nitrogen chemistry. However, according to the  
340 mechanism of this study,  $\text{NO}_2$  only acted as an initiator of the chain reactions in  $\text{SO}_2$  oxidation and essentially all  
341 the  $\text{SO}_2$  was oxidized by  $\text{O}_2$ . Therefore, the amount of HONO formation per sulfate formed was trivial. The  
342 oxidation of  $\text{SO}_2$  by  $\text{O}_2/\text{NO}_2$  is expected to be neither an important source of HONO and OH in the atmosphere  
343 nor to have a significant influence on reactive nitrogen chemistry.

344 In this study, we investigated the reaction of  $\text{SO}_2$  with  $\text{O}_2$  in the presence of  $\text{NO}_2$  at three  $\text{O}_2$  concentrations.  
345 The influence of the  $\text{O}_2$  concentration was shown to be significant. Future experimental results with more  $\text{O}_2$   
346 concentration levels would provide more insights into the reaction mechanism and process.

347 In addition, in the ambient atmosphere, the internal mixing of organics with S(IV) in particles may influence  
348 the S(IV) oxidation rate by  $\text{O}_2$  in the presence of  $\text{NO}_2$ . When organics is abundant in particles, for example during  
349 haze episodes in China, it can react with and thus scavenge radical anion carriers such as  $\text{SO}_5^{\cdot-}$  and  $\text{SO}_4^{\cdot-}$   
350 (Herrmann, 2003; Herrmann et al., 2015; Huie, 1995). Therefore, the presence of internally mixed organics can  
351 reduce the effectivity of the potential radical reaction chain and of S(IV) oxidation, which can undermine the  
352 importance of the oxidation by  $\text{O}_2$  in the presence of  $\text{NO}_2$  in the overall S(IV) oxidation.

### 353 **Acknowledgements**

354 This work was supported by Natural Science Foundation Committee of China (41421064, 21190051,  
355 40490265, 91544000) and Ministry of Science and Technology (Grant No. 2002CB410802).

356 **References**

- 357 Brandt, C., Fabian, I., and Vaneldik, R.: Kinetics and mechanism of the iron(III)-catalyzed autoxidation of  
358 sulfur(IV) oxides in aqueous-solution - evidence for the redox cycling of iron in the presence of oxygen and  
359 modeling of the overall reaction-mechanism, *Inorg. Chem.*, 33, 687-701, 10.1021/ic00082a012, 1994.
- 360 Brandt, C., and Vaneldik, R.: Transition metal-catalyzed oxidation of sulfur (IV) oxides. Atmospheric-relevant  
361 processes and mechanisms, *Chem. Rev.*, 95, 119-190, 10.1021/cr00033a006, 1995.
- 362 Chameides, W. L., and Davis, D. D.: The free-radical chemistry of cloud droplets and its impact upon the  
363 composition of rain, *J. Geophys. Res.-Oceans*, 87, 4863-4877, 10.1029/JC087iC07p04863, 1982.
- 364 Cheng, Y. F., Zheng, G. J., Wei, C., Mu, Q., Zheng, B., Wang, Z. B., Gao, M., Zhang, Q., He, K. B., Carmichael,  
365 G., Poschl, U., and Su, H.: Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events  
366 in China, *Sci. Adv.*, 2, 10.1126/sciadv.1601530, 2016.
- 367 Clifton, C. L., Altstein, N., and Huie, R. E.: Rate-constant for the reaction of NO<sub>2</sub> with sulfur(IV) over the pH  
368 range 5.3-13, *Environ. Sci. Technol.*, 22, 586-589, 10.1021/es00170a018, 1988.
- 369 Davidovits, P., Kolb, C. E., Williams, L. R., Jayne, J. T., and Worsnop, D. R.: Mass accommodation and  
370 chemical reactions at gas-liquid interfaces, *Chem. Rev.*, 106, 1323-1354, 10.1021/cr040366k, 2006.
- 371 Deister, U., and Warneck, P.: Photooxidation of sulfite (SO<sub>3</sub><sup>2-</sup>) in aqueous solution, *J. Phys. Chem.*, 94,  
372 2191-2198, 10.1021/j100368a084, 1990.
- 373 Eriksen, T. E.: pH Effects on the pulse radiolysis of deoxygenated aqueous solutions of sulphur dioxide, *Journal*  
374 *of the Chemical Society, Faraday Transactions 1: Physical Chemistry in Condensed Phases*, 70, 208-215,  
375 10.1039/f19747000208, 1974.
- 376 Gao, M., Carmichael, G. R., Wang, Y., Ji, D., Liu, Z., and Wang, Z.: Improving simulations of sulfate aerosols  
377 during winter haze over Northern China: the impacts of heterogeneous oxidation by NO<sub>2</sub>, *Front. Environ. Sci.*  
378 *Eng.*, 10, 16, 10.1007/s11783-016-0878-2, 2016.
- 379 Guo, H., Weber, R. J., and Nenes, A.: High levels of ammonia do not raise fine particle pH sufficiently to yield  
380 nitrogen oxide-dominated sulfate production, *Sci. Rep.*, 7, 12109, 10.1038/s41598-017-11704-0, 2017.
- 381 Hayon, E., Treinin, A., and Wilf, J.: Electronic spectra, photochemistry, and autoxidation mechanism of the  
382 sulfite-bisulfite-pyrosulfite systems. SO<sub>2</sub><sup>-</sup>, SO<sub>3</sub><sup>-</sup>, SO<sub>4</sub><sup>-</sup>, and SO<sub>5</sub><sup>-</sup> radicals, *J. Am. Chem. Soc.*, 94, 47-57,  
383 10.1021/ja00756a009, 1972.
- 384 He, H., Wang, Y., Ma, Q., Ma, J., Chu, B., Ji, D., Tang, G., Liu, C., Zhang, H., and Hao, J.: Mineral dust and  
385 NO<sub>x</sub> promote the conversion of SO<sub>2</sub> to sulfate in heavy pollution days, *Sci. Rep.*, 4, 10.1038/srep04172, 2014.
- 386 Herrmann, H.: Kinetics of aqueous phase reactions relevant for atmospheric chemistry, *Chem. Rev.*, 103,  
387 4691-4716, 10.1021/cr020658q, 2003.
- 388 Herrmann, H., Schaefer, T., Tilgner, A., Styler, S. A., Weller, C., Teich, M., and Otto, T.: Tropospheric  
389 Aqueous-Phase Chemistry: Kinetics, Mechanisms, and Its Coupling to a Changing Gas Phase, *Chem. Rev.*, 115,  
390 4259-4334, 10.1021/cr500447k, 2015.
- 391 Huang, L., Zhao, Y., Li, H., and Chen, Z.: Kinetics of Heterogeneous Reaction of Sulfur Dioxide on Authentic  
392 Mineral Dust: Effects of Relative Humidity and Hydrogen Peroxide, *Environ. Sci. Technol.*, 49, 10797-10805,  
393 10.1021/acs.est.5b03930, 2015.

394 Huie, R. E., and Neta, P.: Kinetics of one-electron transfer-reactions involving ClO<sub>2</sub> and NO<sub>2</sub>, *J. Phys. Chem.*,  
395 90, 1193-1198, 10.1021/j100278a046, 1986.

396 Huie, R. E., and Neta, P.: Rate constants for some oxidations of S(IV) by radicals in aqueous-solutions, *Atmos.*  
397 *Environ.*, 21, 1743-1747, 10.1016/0004-6981(87)90113-2, 1987.

398 Huie, R. E., Clifton, C. L., and Altstein, N.: A pulse radiolysis and flash photolysis study of the radicals SO<sub>2</sub>,  
399 SO<sub>3</sub>, SO<sub>4</sub> and SO<sub>5</sub>, *Radiat. Phys. Chem.*, 33, 361-370, 1989.

400 Huie, R. E.: Free radical chemistry of the atmospheric aqueous phase, in: *Progress and Problems in Atmospheric*  
401 *Chemistry*, WORLD SCIENTIFIC, 374-419, 1995.

402 Lee, Y.-N., and Schwartz, S. E.: Kinetics of oxidation of aqueous sulfur (IV) by nitrogen dioxide, in:  
403 *Precipitation Scavenging, Dry Deposition and Resuspension*, edited by: Pruppacher, H. R., Semonin, R. G., and  
404 Slinn, W. G. N., Elsevier, New York, 453-466, 1983.

405 Lelieveld, J., Gromov, S., Pozzer, A., and Taraborrelli, D.: Global tropospheric hydroxyl distribution, budget and  
406 reactivity, *Atmos. Chem. Phys.*, 16, 12477-12493, 10.5194/acp-16-12477-2016, 2016.

407 Li, H. J., Zhu, T., Zhao, D. F., Zhang, Z. F., and Chen, Z. M.: Kinetics and mechanisms of heterogeneous  
408 reaction of NO<sub>2</sub> on CaCO<sub>3</sub> surfaces under dry and wet conditions, *Atmos. Chem. Phys.*, 10, 463-474, 2010.

409 Littlejohn, D., Wang, Y. Z., and Chang, S. G.: Oxidation of aqueous sulfite ion by nitrogen-dioxide, *Environ. Sci.*  
410 *Technol.*, 27, 2162-2167, 10.1021/es00047a024, 1993.

411 Liu, M. X., Song, Y., Zhou, T., Xu, Z. Y., Yan, C. Q., Zheng, M., Wu, Z. J., Hu, M., Wu, Y. S., and Zhu, T.: Fine  
412 particle pH during severe haze episodes in northern China, *Geophys. Res. Lett.*, 44, 5213-5221,  
413 10.1002/2017gl073210, 2017.

414 Liu, Y. J., Zhu, T., Zhao, D. F., and Zhang, Z. F.: Investigation of the hygroscopic properties of Ca(NO<sub>3</sub>)(2) and  
415 internally mixed Ca(NO<sub>3</sub>)(2)/CaCO<sub>3</sub> particles by micro-Raman spectrometry, *Atmos. Chem. Phys.*, 8,  
416 7205-7215, 2008.

417 Ma, Q., He, H., Liu, Y., Liu, C., and Grassian, V. H.: Heterogeneous and multiphase formation pathways of  
418 gypsum in the atmosphere, *Phys. Chem. Chem. Phys.*, 15, 19196-19204, 10.1039/c3cp53424c, 2013.

419 Nakamoto, K.: *Infrared and Raman Spectra of Inorganic and Coordination Compounds Part A*, John Wiley  
420 & Sons, New York, 221-247 pp., 1997.

421 Nash, T.: Effect of nitrogen-dioxide and of some transition-metals on the oxidation of dilute bisulfite solutions,  
422 *Atmos. Environ.*, 13, 1149-1154, 10.1016/0004-6981(79)90038-6, 1979.

423 Pöschl, U., Rudich, Y., and Ammann, M.: Kinetic model framework for aerosol and cloud surface chemistry and  
424 gas-particle interactions - Part 1: General equations, parameters, and terminology, *Atmos. Chem. Phys.*, 7,  
425 5989-6023, 2007.

426 Podkrajšek, B., Grgić, I., and Turšič, J.: Determination of sulfur oxides formed during the S(IV) oxidation in the  
427 presence of iron, *Chemosphere*, 49, 271-277, [https://doi.org/10.1016/S0045-6535\(02\)00324-7](https://doi.org/10.1016/S0045-6535(02)00324-7), 2002.

428 Prinn, R. G., Huang, J., Weiss, R. F., Cunnold, D. M., Fraser, P. J., Simmonds, P. G., McCulloch, A., Harth, C.,  
429 Reimann, S., Salameh, P., O'Doherty, S., Wang, R. H. J., Porter, L. W., Miller, B. R., and Krummel, P. B.:  
430 Evidence for variability of atmospheric hydroxyl radicals over the past quarter century, *Geophys. Res. Lett.*, 32,  
431 10.1029/2004gl022228, 2005.

432 Santachiara, G., Prodi, F., and Vivarelli, F.: SO<sub>2</sub> oxidation in monodisperse droplets grown on carbon nuclei in  
433 presence of NO<sub>2</sub>, *J. Aerosol Sci.*, 21, S221-S224, 10.1016/0021-8502(90)90224-1, 1990.

434 Santachiara, G., Prodi, F., and Vivarelli, F.: Further experiments on SO<sub>2</sub> oxidation rate in monodisperse droplets  
435 grown on carbon nuclei in presence of O<sub>2</sub> and NO<sub>2</sub>, *J. Aerosol Sci.*, 24, 683-685,  
436 10.1016/0021-8502(93)90024-4, 1993.

437 Sarma, L. P., Prasad, P. S. R., and Ravikumar, N.: Raman spectroscopic study of phase transitions in natural  
438 gypsum, *J. Raman Spectrosc.*, 29, 851-856, 10.1002/(sici)1097-4555(199809)29:9<851::aid-jrs313>3.0.co;2-s,  
439 1998.

440 Seinfeld, J. H., and Pandis, S. N.: *Atmospheric chemistry and physics: from air pollution to climate change*, 2nd  
441 ed., John Wiley & Sons. Inc., 2006.

442 Shen, C. H., and Rochelle, G. T.: Nitrogen Dioxide Absorption and Sulfite Oxidation in Aqueous Sulfite, *Environ.*  
443 *Sci. Technol.*, 32, 1994-2003, 10.1021/es970466q, 1998.

444 Shi, X. L.: Generation of SO<sub>3</sub><sup>-</sup> and OH radicals in SO<sub>3</sub><sup>2-</sup> reactions with inorganic environmental-pollutants and its  
445 implications to SO<sub>3</sub><sup>2-</sup> toxicity, *J. Inorg. Biochem.*, 56, 155-165, 10.1016/0162-0134(94)85002-x, 1994.

446 Spindler, G., Hesper, J., Brüggemann, E., Dubois, R., Müller, T., and Herrmann, H.: Wet annular denuder  
447 measurements of nitrous acid: laboratory study of the artefact reaction of NO<sub>2</sub> with S(IV) in aqueous solution  
448 and comparison with field measurements, *Atmos. Environ.*, 37, 2643-2662, 10.1016/s1352-2310(03)00209-7,  
449 2003.

450 Turšič, J., Grgić, I., and Bizjak, M.: Influence of NO<sub>2</sub> and dissolved iron on the S(IV) oxidation in synthetic  
451 aqueous solution, *Atmos. Environ.*, 35, 97-104, [https://doi.org/10.1016/S1352-2310\(00\)00283-1](https://doi.org/10.1016/S1352-2310(00)00283-1), 2001.

452 Wang, G., Zhang, R., Gomez, M. E., Yang, L., Levy Zamora, M., Hu, M., Lin, Y., Peng, J., Guo, S., Meng, J., Li,  
453 J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J., Cao, J., An, Z., Zhou, W., Li, G., Wang, J., Tian, P.,  
454 Marrero-Ortiz, W., Secrest, J., Du, Z., Zheng, J., Shang, D., Zeng, L., Shao, M., Wang, W., Huang, Y., Wang, Y.,  
455 Zhu, Y., Li, Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F., Rosenfeld, D., Liss, P. S., Duce, R. A., Kolb,  
456 C. E., and Molina, M. J.: Persistent sulfate formation from London Fog to Chinese haze, *Proc. Nat. Acad. Sci.*  
457 *U.S.A.*, 113, 13630-13635, 10.1073/pnas.1616540113, 2016.

458 Waygood, S. J., and McElroy, W. J.: Spectroscopy and decay kinetics of the sulfite radical anion in aqueous  
459 solution, *J. Chem. Soc.-Faraday Trans.*, 88, 1525-1530, 10.1039/ft9928801525, 1992.

460 Xue, J., Yuan, Z. B., Griffith, S. M., Yu, X., Lau, A. K. H., and Yu, J. Z.: Sulfate Formation Enhanced by a  
461 Cocktail of High NO<sub>x</sub>, SO<sub>2</sub>, Particulate Matter, and Droplet pH during Haze-Fog Events in Megacities in China:  
462 An Observation-Based Modeling Investigation, *Environ. Sci. Technol.*, 50, 7325-7334, 10.1021/acs.est.6b00768,  
463 2016.

464 Zhang, J. K., Sun, Y., Liu, Z. R., Ji, D. S., Hu, B., Liu, Q., and Wang, Y. S.: Characterization of submicron  
465 aerosols during a month of serious pollution in Beijing, 2013, *Atmos. Chem. Phys.*, 14, 2887-2903,  
466 10.5194/acp-14-2887-2014, 2014.

467 Zhao, D., Song, X., Zhu, T., Zhang, Z., and Liu, Y.: Multiphase Reaction of SO<sub>2</sub> with NO<sub>2</sub> on CaCO<sub>3</sub> Particles.  
468 1. Oxidation of SO<sub>2</sub> by NO<sub>2</sub>, *Atmos. Chem. Phys. Discuss.*, 2017, 1-23, 10.5194/acp-2017-610, 2017.

469 Zhao, D. F., Zhu, T., Chen, Q., Liu, Y. J., and Zhang, Z. F.: Raman micro-spectrometry as a technique for  
470 investigating heterogeneous reactions on individual atmospheric particles, *Sci. China Chem.*, 54, 154-160,  
471 10.1007/s11426-010-4182-x, 2011.

472 Zheng, B., Zhang, Q., Zhang, Y., He, K. B., Wang, K., Zheng, G. J., Duan, F. K., Ma, Y. L., and Kimoto, T.:  
473 Heterogeneous chemistry: a mechanism missing in current models to explain secondary inorganic aerosol  
474 formation during the January 2013 haze episode in North China, *Atmos. Chem. Phys.*, 15, 2031-2049,  
475 10.5194/acp-15-2031-2015, 2015a.

476 Zheng, G. J., Duan, F. K., Su, H., Ma, Y. L., Cheng, Y., Zheng, B., Zhang, Q., Huang, T., Kimoto, T., Chang, D.,  
477 Poschl, U., Cheng, Y. F., and He, K. B.: Exploring the severe winter haze in Beijing: the impact of synoptic  
478 weather, regional transport and heterogeneous reactions, *Atmos. Chem. Phys.*, 15, 2969-2983,  
479 10.5194/acp-15-2969-2015, 2015b.

480 Zhu, T., Shang, J., and Zhao, D. F.: The roles of heterogeneous chemical processes in the formation of an air  
481 pollution complex and gray haze, *Sci. China Chem.*, 54, 145-153, 10.1007/s11426-010-4181-y, 2011.

482

483

484

485

486 **Table 1.** Reactive uptake coefficient of SO<sub>2</sub> for sulfate formation at 82% RH and at different O<sub>2</sub>  
487 concentrations.

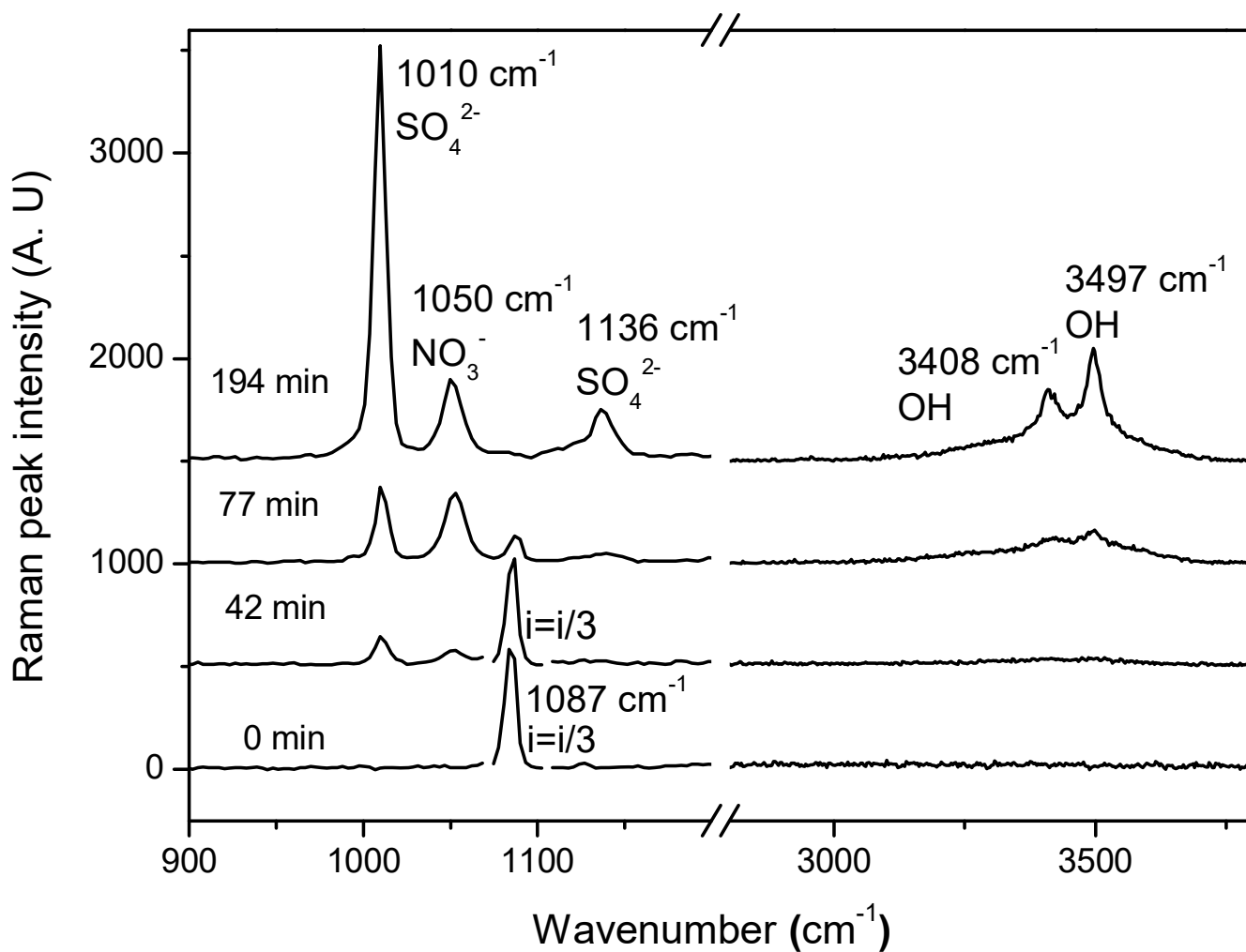
SO <sub>2</sub> /NO <sub>2</sub> /O <sub>2</sub> concentration	$\gamma$
75 ppm/ 75 ppm/ 86 %	$1.7 \times 10^{-5}$
75 ppm/ 75 ppm/ 20 %	$1.2 \times 10^{-5}$
75 ppm/ 75 ppm/ 5 %	$3.5 \times 10^{-6}$

488

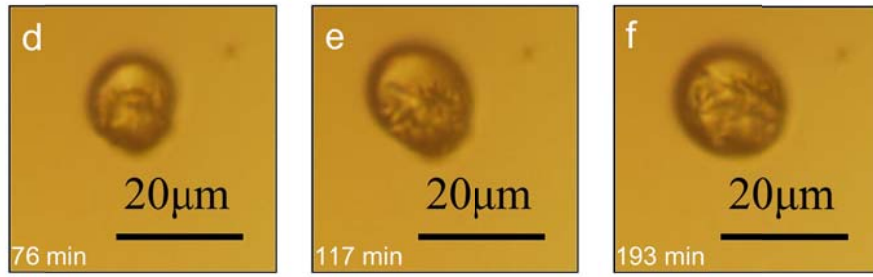
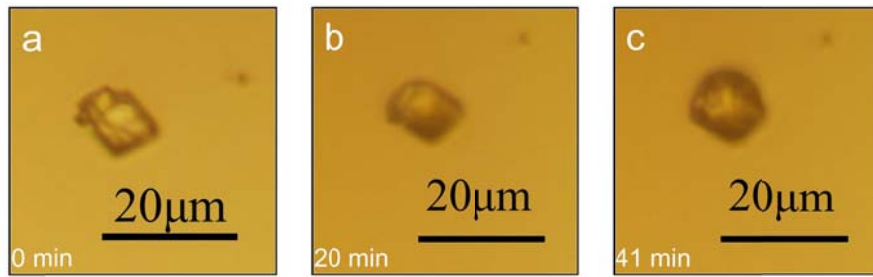
**Table 2.** Summary of the mechanism of the reaction S(IV) with O<sub>2</sub>/NO<sub>2</sub>

Step	Reactions
Initiation	$\text{NO}_2(\text{aq}) + \text{HSO}_3^-(\text{aq}) \rightarrow \text{NO}_2^-(\text{aq}) + \text{SO}_3^{\bullet-}(\text{aq}) + \text{H}^+(\text{aq})$ (R8)
Propagation	$\text{SO}_3^{\bullet-}(\text{aq}) + \text{O}_2(\text{aq}) \rightarrow \text{SO}_5^{\bullet-}(\text{aq})$ (R12)
	$\text{SO}_5^{\bullet-}(\text{aq}) + \text{HSO}_3^-(\text{aq}) \rightarrow \text{HSO}_5^-(\text{aq}) + \text{SO}_3^{\bullet-}(\text{aq})$ (R13)
	$\text{HSO}_5^-(\text{aq}) + \text{HSO}_3^-(\text{aq}) \rightarrow 2\text{SO}_4^{2-}(\text{aq}) + 2\text{H}^+(\text{aq})$ (R14)
	$\text{SO}_5^{\bullet-}(\text{aq}) + \text{HSO}_3^-(\text{aq}) \rightarrow \text{SO}_4^{2-}(\text{aq}) + \text{SO}_4^{\bullet-}(\text{aq}) + \text{H}^+(\text{aq})$ (R15)
	$\text{SO}_4^{\bullet-}(\text{aq}) + \text{HSO}_3^-(\text{aq}) \rightarrow \text{SO}_4^{2-}(\text{aq}) + \text{SO}_3^{\bullet-}(\text{aq}) + \text{H}^+(\text{aq})$ (R16)
Termination	$\text{SO}_3^{\bullet-}(\text{aq}) + \text{SO}_3^{\bullet-}(\text{aq}) \rightarrow \text{S}_2\text{O}_6^{2-}(\text{aq})$ (R9)
	$\text{SO}_3^{\bullet-}(\text{aq}) + \text{SO}_3^{\bullet-}(\text{aq}) \rightarrow \text{SO}_3^{2-}(\text{aq}) + \text{SO}_3$ (R10)
	$\text{SO}_3(\text{aq}) + \text{H}_2\text{O} \rightarrow \text{SO}_4^{2-}(\text{aq}) + 2\text{H}^+(\text{aq})$ (R11)
	$\text{SO}_4^{\bullet-}(\text{aq}) + \text{SO}_4^{\bullet-}(\text{aq}) \rightarrow \text{S}_2\text{O}_8^{2-}(\text{aq})$ (R17)
	$\text{SO}_5^{\bullet-}(\text{aq}) + \text{SO}_5^{\bullet-}(\text{aq}) \rightarrow \text{S}_2\text{O}_8^{2-}(\text{aq}) + \text{O}_2(\text{aq})$ (R18)

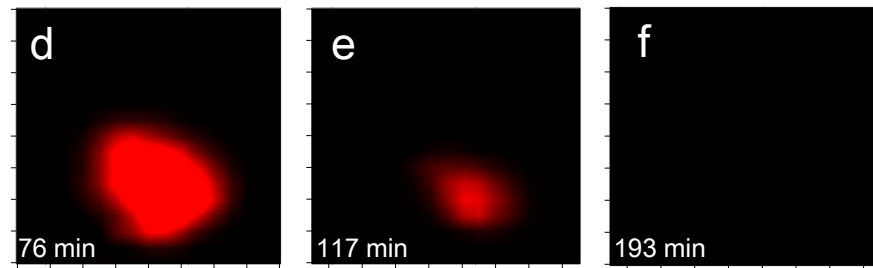
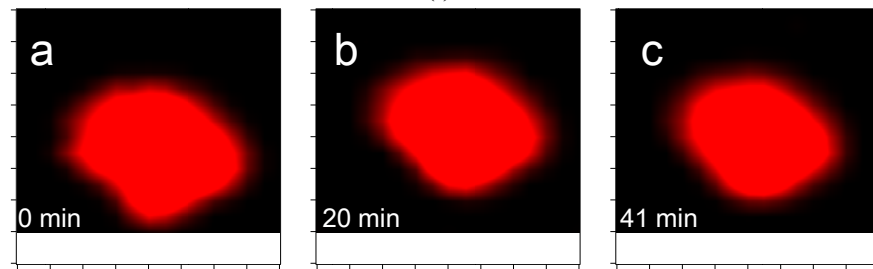




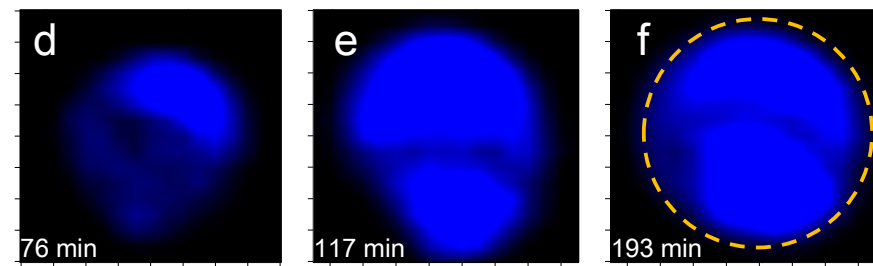
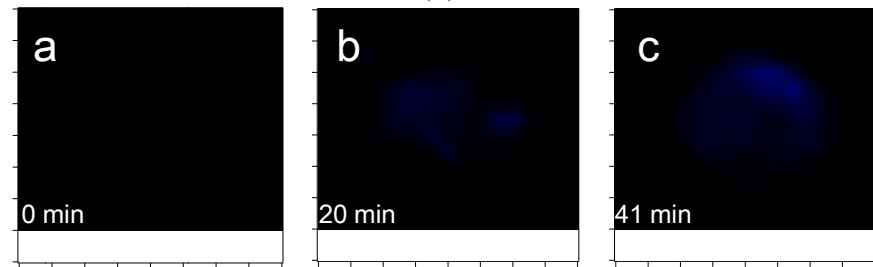
491  
 492 Figure 1. Raman spectra of a  $\text{CaCO}_3$  particle during the multiphase reaction of  $\text{SO}_2$  with  $\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$   
 493 on the particle.  $\text{SO}_2$ : 75 ppm,  $\text{NO}_2$ : 75 ppm, RH: 72%,  $\text{O}_2$ : 20%. The peak intensity of carbonate ( $1087$   
 494  $\text{cm}^{-1}$ ) at 0 and 42 min was divided by three for clarity.  
 495



(i)



(ii)

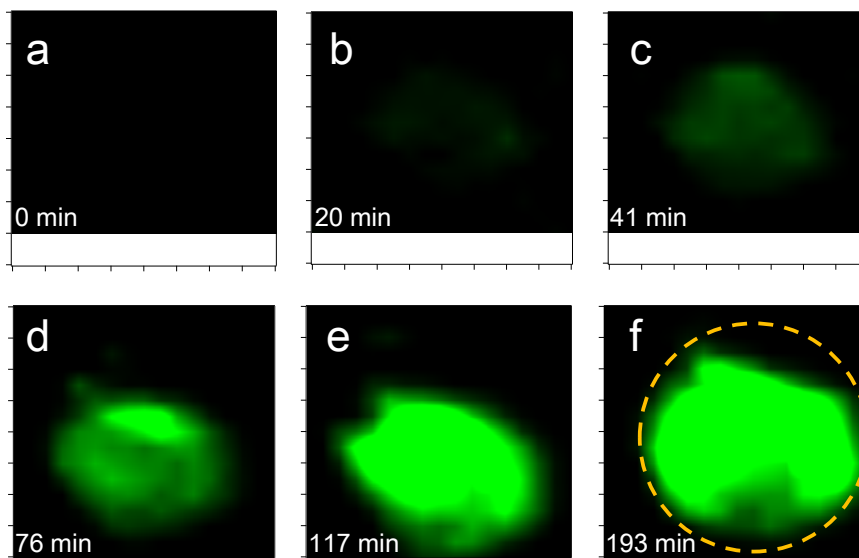


(iii)

496  
497

498  
499

500  
501



502

503

504 (iv) Figure 2. Microscopic image (i) and Raman mapping image of carbonate (ii), nitrate (iii), and sulfate (iv)

505 on the  $\text{CaCO}_3$  particle during the multiphase reaction  $\text{SO}_2$  with  $\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$  on the particle. A-f

506 corresponds to the reaction time of 0, 20, 41, 76, 117, and 193 min.  $\text{SO}_2$ : 75 ppm,  $\text{NO}_2$ : 75 ppm, RH:

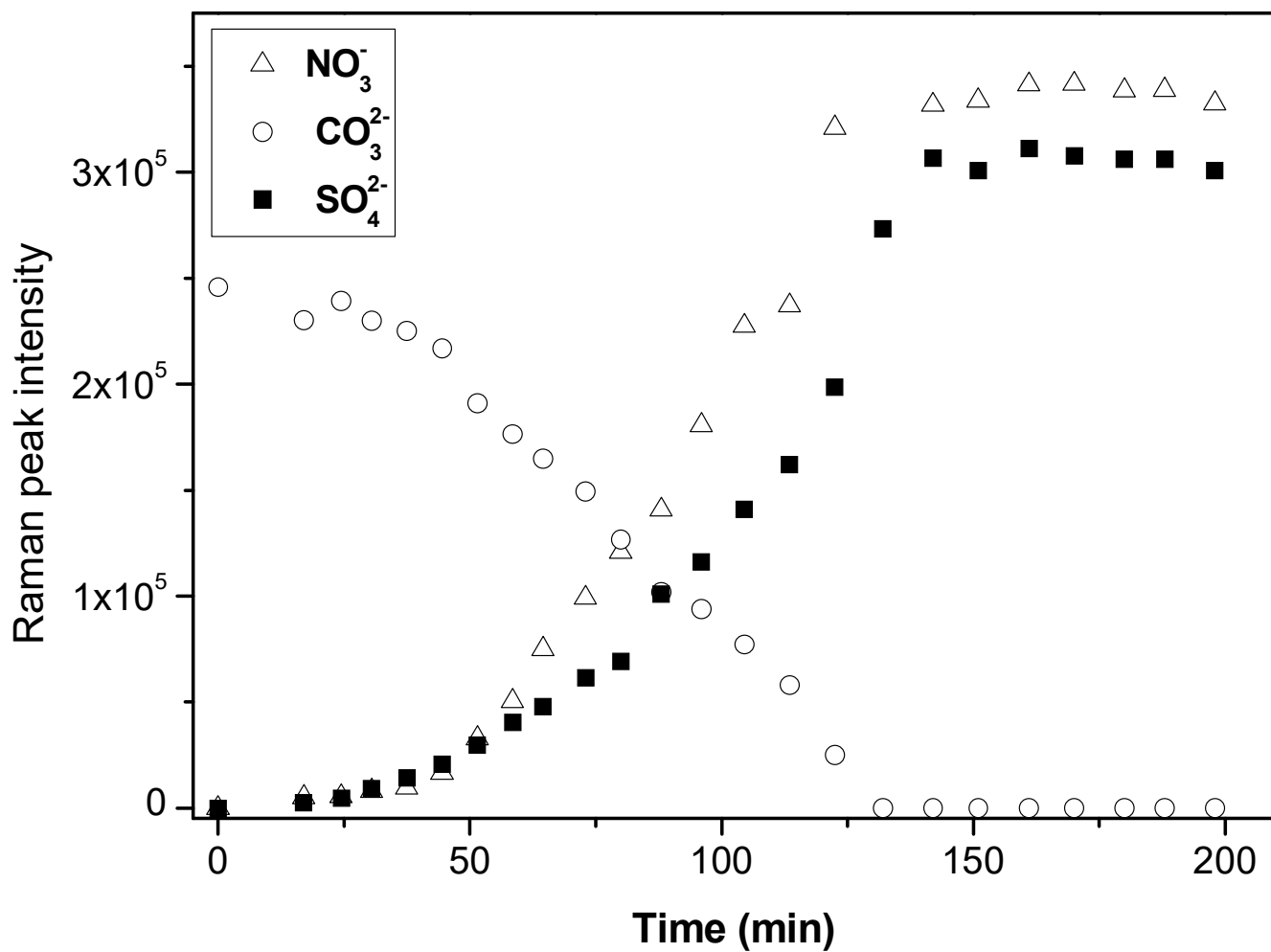
507 72%,  $\text{O}_2$ : 20%. The mapping image of carbonate, nitrate, and sulfate are made using the peak area at

508 1050, 1010, and 1087  $\text{cm}^{-1}$ , respectively. The red, blue, and green colors indicate the peak intensity of

509 carbonate, nitrate, and sulfate, respectively. The dashed lines in panel iii-f and iv-f indicate the shape of

510 the droplet at the end of the reaction.

511



512  
 513 Figure 3. Time series of the Raman peak intensity of  $\text{NO}_3^-$ ,  $\text{SO}_4^{2-}$ , and  $\text{CO}_3^{2-}$  during the reaction of  $\text{SO}_2$   
 514 with  $\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$  on  $\text{CaCO}_3$  particles.  $\text{SO}_2$ : 75 ppm,  $\text{NO}_2$ : 75 ppm, RH: 72%,  $\text{O}_2$ : 20%. The intensity of  
 515  $\text{NO}_3^-$ ,  $\text{SO}_4^{2-}$ , and  $\text{CO}_3^{2-}$  show the peak area at 1050, 1010, and 1087  $\text{cm}^{-1}$ , respectively, in Raman spectra  
 516 obtained by Raman mapping.

517  
 518