

Response to referee #1

We thank the reviewer for further reviewing our manuscript and providing helpful comments. We have addressed these comments. In the following please find our response and the corresponding change made to the manuscript. The original comments are shown in italics. The changes made in the revised manuscript are highlighted.

I feel that the authors have addressed my most serious concerns from the previous version. However, before the article is suitable for publication, the authors need to address the fact that these experiments were conducted at extremely high, environmentally unrealistic reactant concentrations (75 ppm SO₂ and NO₂). This is particularly relevant since the reaction products are expected to alter the properties of the particle surface (pH, hygroscopicity) and this process will occur unrealistically fast at elevated reactant concentrations.

Response:

We have accepted the suggestion. In our experiments, the SO₂ and NO₂ concentrations much higher than those in ambient levels were used, because 1) enough sulfate as a reaction product can be formed on a single particle to obtain high signal-to-noise ratio in our experimental setup; 2) in laboratory studies, high concentrations of reactants are often used in order to simulate the chemical/physical processes within minutes or hours which occur in the ambient atmosphere on the time scale of days or weeks.

In the revised manuscript, we have discussed this limitation as follows.

“It is important to note that the concentrations of NO₂ and SO₂ used in this study are much higher than those in the ambient atmosphere. High concentrations of reactant gases are often used in laboratory studies in order to simulate the ambient reactions at the time scale of days or weeks and to get high signal-to-noise ratios for detecting products within minutes or hours. In the ambient atmosphere, reactive uptake coefficient of SO₂ should be lower than that in this study due to the lower NO₂ concentrations when other conditions are comparable and the chemical/physical processes observed in this study, such as changes in particle composition, phase, hygroscopicity, and pH should be much slower due to the lower concentrations of NO₂ and SO₂.”

Response to referee #2

We thank the reviewer for carefully reviewing our manuscript and providing helpful comments. All the comments have been addressed in the revised manuscript and we believe that the revisions based on these comments have improved our manuscript significantly. In the following please find our responses to the comments one by one, and the corresponding changes made to the manuscript. The original comments are shown in italics. The changes made in the revised manuscript are highlighted.

General: I am receiving the revised manuscript without having reviewed it before. I think the paper is generally ok and the authors have addressed the concerns of the reviewers.

The authors tend to see the observed oxygen-dependence as a support for the direct single electron transfer between HSO_3^- and NO_2 which directly leads to SO_3^- and $\text{NO}_2^- / \text{HONO}$. The finding of the oxygen effect is interesting but possible more measured points would have been helpful – that could be included into an outlook. Please also note that the NO_2 and SO_2 mixing ratios applied are gigantic. Regarding the mechanism of the reaction between S(IV) and N(IV) thermochemistry should be considered. Such treatment shows that the reaction is only exothermic by a very small extend and it should be slow. The complex formation mechanism appears much more feasible to me because of this. Maybe oxygen can promote the decomposition of the adduct by scavenging the formed SO_3^- , if there is not only a unimolecular decomposition of the complex but an equilibrium. Then oxygen would be able to promote scavenging of SO_3^- to SO_5^- thus shifting a decomposition equilibrium on the product side. Maybe this could be discussed and included into the discussion / outlook.

Anyway, whether there is a concerted single electron transfer or whether the reaction runs through a complex, at the end sulfur-oxy radical anions will be produced the role of which is S(IV) oxidation has been studied extensively in the late 1980s and early 1990 both in the US and in Europe. These reactions are included into state-of-the art aqueous chemical mechanism and a chemistry such as suggested here would add a source of SO_3^- . One thought which is important for the whole scope of the paper: Such an additional SO_3^- radical anion source will not lead to a much increased S(VI) production as the sulfur radical chain will not establish but the reactive chain carriers will be scavenged and reduce the effectivity of the potential radical reaction chain. This has been demonstrate in many studies and , to some extent, it is surprising that this is not being discussed adequately in a study motivated to explain particle sulfate formation correlating to NO_2 in the gas phase. Especially in China, the sulfur-oxy radical chemistry will be scavenged by the abundance of

organics in particles – this occurs via the reaction of SO_4^- with organics where many kinetic data are available but also by such reactions of SO_5^- where the data basis is more sparse. This overall thought should surely be considered in the study.

Response:

We thank the reviewer for supportive remarks.

The effect of oxygen was shown to be significant. We agree with the reviewer that more oxygen levels would be helpful. In the revised manuscript, we have added a brief outlook to have experimental results with more oxygen levels as follows.

“In this study, we investigated the reaction of SO_2 with O_2 in the presence of NO_2 at three O_2 concentrations. The influence of the O_2 concentration was shown to be significant. Future experimental results with more O_2 concentration levels would provide more insights into the reaction mechanism and process.”

Moreover, we have discussed the limitation of using high mixing ratios of SO_2 and NO_2 in the revised manuscript as follows.

“It is important to note that the concentrations of NO_2 and SO_2 used in this study are much higher than those in the ambient atmosphere. High concentrations of reactant gases are often used in laboratory studies in order to simulate the ambient reactions at the time scale of days or weeks and to get high signal-to-noise ratios for detecting products within minutes or hours. In the ambient atmosphere, reactive uptake coefficient of SO_2 should be lower than that in this study due to the lower NO_2 concentrations when other conditions are comparable and the chemical/physical processes observed in this study, such as changes in particle composition, phase, hygroscopicity, and pH should be much slower due to the lower concentrations of NO_2 and SO_2 .”

We agree that the mechanism involving the decomposition of the reaction complex and the equilibrium as the reviewer suggested is possible. In the revised manuscript, we have added discussion on this mechanism as follows.

“In addition to the two mechanisms above, Spindler et al. (2003) proposed a reaction mechanism involving first NO_2 -S(IV) complex formation and subsequent $SO_3^{\cdot-}$ radical formation (R3, R7). NO_2 -S(IV) complex may establish an equilibrium with $SO_3^{\cdot-}$ in contrast to the direct formation of $SO_3^{\cdot-}$ via the reaction of NO_2 with SO_2 . Higher concentration of O_2 favors the conversion of $SO_3^{\cdot-}$ to $SO_5^{\cdot-}$ and thus can reduce the $SO_3^{\cdot-}$ concentration, shifting the equilibrium to the product side and promoting the

overall S(IV) oxidation. O₂ can act in a similar way as in the free-radical chain mechanism. Admittedly, we cannot rule out the possibility NO₂-S(IV) complex formation. But such a mechanism can still be classified as the free-radical chain mechanism since the S(IV) oxidation still proceeds via the radical chain reactions.”

We agree that in the ambient atmosphere in the internally mixed particles where organics and S(IV) co-exist, organics in particles can scavenge sulfur-oxy radical anions and thus reduce the effectivity of the potential radical reaction chain and of S(IV) oxidation. In the revised manuscript, we have discussed this point as follows.

“In addition, in the ambient atmosphere, the internal mixing of organics with S(IV) in particles may influence the S(IV) oxidation rate by O₂ in the presence of NO₂. When organics is abundant in particles, for example during haze episodes in China, it can react with and thus scavenge radical anion carriers such as SO₅⁻ and SO₄⁻ (Herrmann, 2003; Herrmann et al., 2015; Huie, 1995). Therefore, the presence of internally mixed organics can reduce the effectivity of the potential radical reaction chain and of S(IV) oxidation, which can undermine the importance of the oxidation by O₂ in the presence of NO₂ in the overall S(IV) oxidation.”

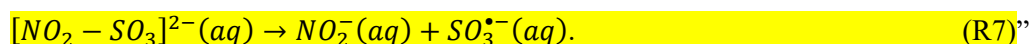
Details

Line 53: The postulation of complexes involving more than one NO₂ unit probably comes from pulse radiolysis experiments with quite high NO₂ concentrations. It should be considered that not only the complex with three NO₂ unit can decompose but also the others, especially the most simple one being formed in R3.

Response:

Accepted. In the revised manuscript, we have added the reaction equation of the decomposition of simple SO₂-NO₂ complex.

“Additionally, NO₂-S(IV) adduct complex may decompose as follows (Spindler et al., 2003).



L 205-208: See above remarks, I would be a bit more cautious here.

Response:

Accepted. In the revised manuscript, we have revised this part.

“According to the NO₂-S(IV) complex mechanism, the presence of O₂ should not affect the SO₂ oxidation rate; however, in this study, a substantial enhancement in the SO₂ oxidation rate was observed in the presence of O₂ compared with that in the absence of O₂. Therefore, the NO₂-S(IV) complex mechanism was less likely to have been important in this study.”

We further discussed the mechanism involving both NO₂-S(IV) complex and SO₃^{•-} radical as mentioned above.

“In addition to the two mechanisms above, Spindler et al. (2003) proposed a reaction mechanism involving first NO₂-S(IV) complex formation and subsequent SO₃^{•-} radical formation (R3, R7). NO₂-S(IV) complex may establish an equilibrium with SO₃^{•-} in contrast to the direct formation of SO₃^{•-} via the reaction of NO₂ with SO₂. Higher concentration of O₂ favors the conversion of SO₃^{•-} to SO₅^{•-} and thus can reduce the SO₃^{•-} concentration, shifting the equilibrium to the product side and promoting the overall S(IV) oxidation. O₂ can act in a similar way as in the free-radical chain mechanism. Admittedly, we cannot rule out the possibility NO₂-S(IV) complex formation. But such a mechanism can still be classified as the free-radical chain mechanism since the S(IV) oxidation still proceeds via the radical chain reactions.”

L 209-257: All of this must be put into context with the sensitivity of the sulfur-oxy-radical reaction chain towards organics under real environmental conditions.

Response:

Accepted. In the revised manuscript, we have discussed the influence of organics in the particle phase on the sulfur-oxy-radical chain reactions as in the response to the general comments above.

L 286 ff: Conclusions section: I feel this should be modified somewhat in view of the above.

Response:

Accepted. In the revised manuscript, we have modified this section by adding the following text.

“In this study, we investigated the reaction of SO₂ with O₂ in the presence of NO₂ at three O₂ concentrations. The influence of the O₂ concentration was shown to be significant. Future experimental results with more O₂ concentration levels would provide more insights into the reaction mechanism and process.”

In addition, in the ambient atmosphere, the internal mixing of organics with S(IV) in particles may influence the S(IV) oxidation rate by O_2 in the presence of NO_2 . When organics is abundant in particles, for example during haze episodes in China, it can react with and thus scavenge radical anion carriers such as $SO_5^{\cdot-}$ and $SO_4^{\cdot-}$ (Herrmann, 2003; Herrmann et al., 2015; Huie, 1995). Therefore, the presence of internally mixed organics can reduce the effectivity of the potential radical reaction chain and of S(IV) oxidation, which can undermine the importance of the oxidation by O_2 in the presence of NO_2 in the overall S(IV) oxidation.”

1 **NO₂-initiated multiphase oxidation of SO₂ by O₂ on CaCO₃ particles**

2 Ting Yu*, Defeng Zhao*, Xiaojuan Song, Tong Zhu

3 BIC-ESAT and SKL-ESPC, College of Environmental Sciences and Engineering, Peking University, Beijing,
4 100871, China

5 *These authors contributed equally to this work.

6 *Correspondence to:* Tong Zhu (tzhu@pku.edu.cn)

7 **Abstract.** The reaction of SO₂ with NO₂ on the surface of aerosol particles has been suggested to be important in
8 sulfate formation during severe air pollution episodes in China. However, we found that the direct oxidation of
9 SO₂ by NO₂ was slow and might not be the main reason for sulfate formation in ambient air. In this study, we
10 investigated the multiphase reaction of SO₂ with an O₂/NO₂ mixture on single CaCO₃ particles using
11 Micro-Raman spectroscopy. The reaction converted the CaCO₃ particle to a Ca(NO₃)₂ droplet, with CaSO₄•2H₂O
12 solid particles embedded in it, which constituted a significant fraction of the droplet volume at the end of the
13 reaction. The reactive uptake coefficient of SO₂ for sulfate formation was on the order of 10⁻⁵, which was higher
14 than that for the multiphase reaction of SO₂ directly with NO₂ by 2–3 orders of magnitude. According to our
15 observations and the literature, we found that in the multiphase reaction of SO₂ with the O₂/NO₂ mixture, O₂ was
16 the main oxidant of SO₂ and was necessary for radical chain propagation. NO₂ acted as the initiator of radical
17 formation, but not as the main oxidant. The synergy of NO₂ and O₂ resulted in much faster sulfate formation than
18 the sum of the reaction rates with NO₂ and with O₂ alone. We estimated that the multiphase oxidation of SO₂ by
19 O₂ initiated by NO₂ could be an important source of sulfate and a sink of SO₂, based on the calculated lifetime of
20 SO₂ regarding the loss through the multiphase reaction versus the loss through the gas-phase reaction with OH
21 radical. Parameterizing the reactive uptake coefficient of the reaction observed in our laboratory for further model
22 simulation is needed, as well as an integrated assessment based on field observations, laboratory study results,
23 and model simulations to evaluate the importance of the reaction in ambient air during severe air pollution
24 episodes, especially in China.

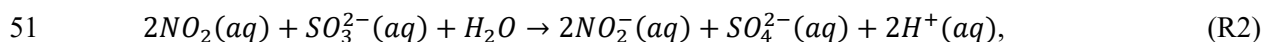
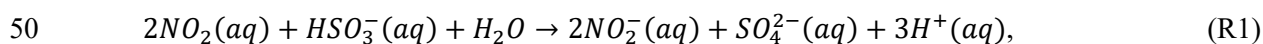
25 1 Introduction

26 It has been suggested that multiphase or heterogeneous oxidation of SO₂ potentially plays an important role
27 in sulfate formation in the atmosphere (Seinfeld and Pandis, 2006). During the severe pollution episodes that
28 occur frequently in China, high sulfate concentrations cannot be explained by the gas phase oxidation of SO₂ and
29 its well-known aqueous chemistry (Zheng et al., 2015a; Cheng et al., 2016), highlighting the role of
30 under-appreciated heterogeneous oxidation or multiphase pathways.

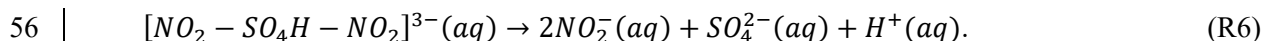
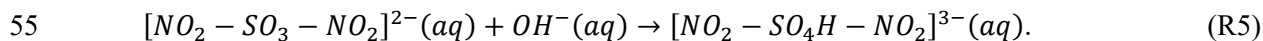
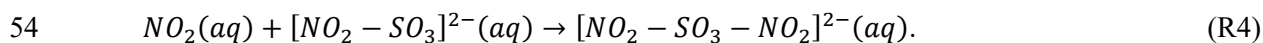
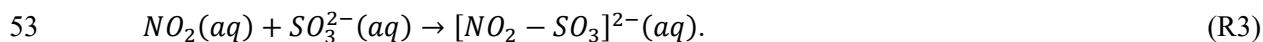
31 Recently, the multiphase oxidation of SO₂ by NO₂ has been introduced in air quality model simulations to
32 explain the discrepancy between the modeled and observed sulfate concentration during severe pollution episodes
33 in China (Cheng et al., 2016; Gao et al., 2016; Wang et al., 2016; Xue et al., 2016), despite the uncertainties in
34 the kinetic parameters for SO₂ oxidation and in the pH value of aerosol particles in China (Wang et al., 2016;
35 Cheng et al., 2016; Liu et al., 2017; Guo et al., 2017). However, according to our recently published results (Zhao
36 et al., 2017), the direct oxidation of SO₂ by NO₂ could not contribute significantly to sulfate formation in the
37 atmosphere because the reactive uptake coefficient of SO₂ for sulfate formation due to direct oxidation by NO₂ is
38 very low (~10⁻⁸).

39 Although the contribution of the direct oxidation of SO₂ by NO₂ to sulfate formation is not significant, NO₂
40 may be involved in other oxidation pathways of SO₂. It has been reported that the reaction of NO₂ with SO₃²⁻ and
41 HSO₃⁻ in the bulk aqueous phase can form the SO₃^{•-} radical, which can further react with O₂ and produce a series
42 of radicals that oxidize S(IV) species (Littlejohn et al., 1993). The reaction pathway may result in a fast SO₂
43 oxidation due to the potential synergy of NO₂ and O₂.

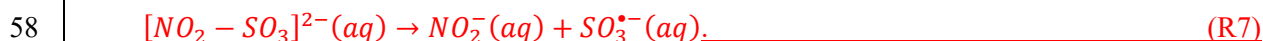
44 Despite such a reaction mechanism for SO₂ oxidation being proposed, its role in SO₂ oxidation in the
45 ambient atmosphere is not well established. Most previous studies have focused on the direct reaction of SO₂
46 with NO₂, including the determination of its rate constant (Lee and Schwartz, 1983; Clifton et al., 1988; Shen and
47 Rochelle, 1998; Spindler et al., 2003; Nash, 1979; Huie and Neta, 1986). According to the reaction products and
48 their reported yields (Lee and Schwartz, 1983; Clifton et al., 1988), the overall reaction equations of the direct
49 reaction of SO₂ with NO₂ are as follows:



52 and the reactions are proposed to proceed via NO₂-S(IV) adduct complexes (Clifton et al., 1988).



57 Additionally, NO₂-S(IV) adduct complex may decompose as follows (Spindler et al., 2003).



59 However, studies of the oxidation rate of SO₂ at the O₂ concentrations relevant to the ambient atmosphere and the
60 potential influence of the synergy of NO₂ and O₂ on the oxidation rate are very limited (Turšič et al., 2001; He et
61 al., 2014), except a few studies investigated SO₂ oxidation in the presence of NO₂ as well as O₂ (Littlejohn et al.,
62 1993; Shen and Rochelle, 1998; Santachiara et al., 1990). Moreover, previous studies have mainly focused on the

63 reaction in bulk solution and only few studies have investigated the oxidation of SO₂ by NO₂ on aerosol particles
64 (Santachiara et al., 1990, 1993). On aerosol particles, water activity, pH, ionic strength, the presence of other
65 compounds or ions, and the role of particle surface are different from in dilute bulk solution and may affect the
66 reaction process and reaction rate. Therefore, further studies of the multiphase reaction of SO₂ with O₂/NO₂
67 mixtures on aerosol particles are required to determine the kinetic parameters and the mechanism of the reaction.

68 In this study, we investigated the multiphase reaction of SO₂ with O₂ in the presence of NO₂ on CaCO₃
69 particles. We quantified the reactive uptake coefficient of SO₂ due to the reaction with an O₂/NO₂/H₂O mixture.
70 Based on our observations and the existing literature, we further discussed the reaction mechanism. Furthermore,
71 we estimated the role of the multiphase oxidation of SO₂ by O₂ in the presence of NO₂ in the atmosphere.

72 2 Experimental

73 The experiments were conducted using a flow reaction system and the setup is shown in Fig. S1. The
74 experimental setup and procedure used have been described in detail in previous studies (Zhao et al., 2017; Zhao
75 et al., 2011; Liu et al., 2008). A gas mixture of NO₂, SO₂, O₂, N₂, and water vapor reacted with particles
76 deposited on a substrate in the flow reaction cell. The concentrations of SO₂ and NO₂ were controlled using mass
77 flow controllers by varying the flow rates of SO₂ (2,000 ppm in high purity N₂, National Institute of Metrology
78 P.R. China), NO₂ (1,000 ppm in high purity N₂, Messer, Germany), and synthetic air [20% O₂ (high purity grade:
79 99.999%, Beijing Haikeyuanchang Practical Gas Co., Ltd.) and 80% N₂ (high purity grade: 99.999%, Beijing
80 Haikeyuanchang Practical Gas Co., Ltd.)]. Relative humidity (RH) was controlled by regulating the flow rates of
81 reactant gases, dry synthetic air, and humidified synthetic air. Humidified synthetic air was prepared by bubbling
82 synthetic air through fritted glass in water. In some experiments, the O₂ concentrations were varied by regulating
83 the mixing ratios of O₂ and N₂ to investigate the effect of O₂. SO₂/O₂/NO₂/H₂O mixtures flew through the
84 reaction cell and reacted with individual stationary CaCO₃ particles, which were deposited on a Teflon-FEP film
85 substrate annealed to a silicon wafer. RH and temperature were measured using a hygrometer (HMT100, Vaisala,
86 Vantaa, Finland) at the exit of the reaction cell. Additionally, temperature was measured using another small
87 temperature sensor (Pt 100, 1/3 DIN B, Heraeus, Hanau, Germany) in the reaction cell. All the experiments were
88 conducted at 298 ± 0.5 K. The experiments were conducted under two RHs (72% and 82%) at 75 ppm SO₂ and
89 75 ppm NO₂.

90 During the reaction, particles were monitored *in-situ* via a glass window on the top of the reaction cell using
91 a Micro-Raman spectrometer (LabRam HR800, HORIBA Jobin Yvon, Kyoto, Japan) to obtain microscopic
92 images and Raman spectra. A 514-nm excitation laser was used, and back scattering Raman signals were detected.
93 The details of the instrument are described elsewhere (Liu et al., 2008; Zhao et al., 2011). Because the particles
94 were larger than the laser spot in this study (~1.5 μm), confocal Raman mapping was used to measure the spectra
95 at different locations on a particle to obtain the chemical information of the entire particle. The mapping area was
96 rectangular and was slightly larger than the particle, with mapping steps of 1 × 1 μm. Raman spectra in the range
97 of 800–3,900 cm⁻¹ were acquired with an exposure time of 1 s for each mapping point. Raman spectra were
98 analyzed using Labspec 5 software (HORIBA Jobin Yvon). Raman peaks were fitted to Gaussian–Lorentzian

99 functions to obtain peak positions and peak areas at different locations on the particle. The peak areas were then
100 added together to obtain the peak area for the entire particle.

101 Particles of CaCO₃ (98%, Sigma-Aldrich, USA), with average diameters of about 7–10 μm as specified by
102 the supplier, were used in the experiments. The CaCO₃ particles were rhombohedron crystals; X-ray diffraction
103 analysis indicated that they were calcite (Fig. S2). Individual particles were prepared by dripping a dilute CaCO₃
104 suspended solution onto Teflon-FEP film using a pipette and then drying the sample in an oven at 80°C for 10 h.

105 The amount of CaSO₄ as a reaction product was quantified based on Raman peak areas and particle sizes.
106 The details of the method are described in our previous study (Zhao et al., 2017). Briefly, the amount of reaction
107 product CaSO₄ formed was determined as a function of time using Raman peak areas. Raman peak areas were
108 converted to the amount of compound formed using a calibration curve obtained from pure CaSO₄ particles of
109 different sizes, which were determined according to microscopic images. The reaction rate, i.e., the sulfate
110 production rate, was derived from the amount of sulfate formed as a function of time. The reactive uptake
111 coefficient of SO₂ for sulfate formation (γ) was further determined from the reaction rate and collision rate of SO₂
112 on the surface of a single particle.

$$113 \quad \gamma = \frac{d\{SO_4^{2-}\}}{Z} \cdot dt \quad (1)$$

$$114 \quad Z = \frac{1}{4}cA_s[SO_2], \quad (2)$$

$$115 \quad c = \sqrt{\frac{8RT}{\pi M_{SO_2}}}, \quad (3)$$

116 where R is the gas constant, T is temperature, M_{SO₂} is the molecular weight of SO₂, c is the mean molecular
117 velocity of SO₂, A_s is the surface area of an individual particle, and Z is the collision rate of SO₂ on the surface of
118 a particle. {SO₄²⁻} indicates the amount of sulfate in the particle phase in moles. The average reaction rate and
119 surface area of particles during the multiphase reaction period were used to derive the reactive uptake coefficient.
120 The period was chosen to start after the induction period when ~10% of the final sulfate was formed. [SO₂]
121 indicates the concentration of SO₂ in the gas phase.

122 The influence of gas phase diffusion on reactive uptake was evaluated using the resistor model described by
123 Davidovits et al. (2006) and references therein, as well as using the gas phase diffusion correction factor for a
124 reactive uptake coefficient according to the method described by Pöschl et al. (2007). The reactive uptake of SO₂
125 was found to not be limited by gas phase diffusion (see details in the Supplement S1).

126 In addition, we conducted experiments of the reaction SO₂ with only O₂ on both CaCO₃ solid particles and
127 internally mixed CaCO₃/Ca(NO₃)₂ particles (with CaCO₃ embedded in Ca(NO₃)₂ droplets), while keeping other
128 conditions the same as the reaction of SO₂ with an O₂/NO₂ mixture. These experiments of the multiphase
129 oxidation of SO₂ by O₂ can help determine the role of NO₂ in the reaction of SO₂ with an O₂/NO₂ mixture.

130 3 Results and discussion

131 3.1 Reaction products and changes in particle morphology

132 Figure 1 shows the Raman spectra of a CaCO₃ particle during the multiphase reaction of SO₂ with O₂/NO₂/H₂O
133 on its surface. The peak at 1,087 cm⁻¹ was assigned to the symmetric stretching of carbonate ($\nu_s(\text{CO}_3^{2-})$)

134 (Nakamoto, 1997). During the reaction, the peak at $1,087\text{ cm}^{-1}$ decreased continuously and finally disappeared as
135 new peaks were observed. The peak at $1,050\text{ cm}^{-1}$ was assigned to the symmetric stretching of nitrate ($\nu_s(\text{NO}_3^-)$).
136 The peaks at $1,010\text{ cm}^{-1}$ and $1,136\text{ cm}^{-1}$ were assigned to the symmetric stretching ($\nu_s(\text{SO}_4^{2-})$) and asymmetric
137 stretching ($\nu_{as}(\text{SO}_4^{2-})$) of sulfate in gypsum ($\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$), respectively (Sarma et al., 1998). In addition, after the
138 reaction, a broad envelope in the range of $2,800\text{--}3,800\text{ cm}^{-1}$ assigned to the stretching of the OH bond in water
139 molecules was observed. Above this envelope, there were two peaks at $3,408\text{ cm}^{-1}$ and $3,497\text{ cm}^{-1}$, which were
140 assigned to OH bond stretching in crystallization water of $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$ (Sarma et al., 1998; Ma et al., 2013).

141 During the multiphase reaction with the $\text{SO}_2/\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$ mixture, the CaCO_3 particles displayed a remarkable
142 change in morphology. The original CaCO_3 particle was a rhombohedron crystal (Fig. 2, panel i, a). As the
143 reaction proceeded, its edges became smoother and later a transparent droplet layer formed, which had a newly
144 formed solid phase embed in it (Fig. 2, panel i, d). The size of the new solid phase grew during the reaction (Fig.
145 2, panel i, d–f) and it seemed to contain many micro-crystals. Raman mapping revealed that the new solid phase
146 consisted of $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$ (Fig. 2, panel iv), and the surrounding aqueous layer consisted of $\text{Ca}(\text{NO}_3)_2$ (Fig. 2,
147 panel iii).

148 The particle morphology change shown in Fig. 2 was significantly different from the morphology change in
149 the direct reaction of SO_2 with NO_2 (Zhao et al., 2017), where the CaCO_3 particle was first converted to a
150 spherical $\text{Ca}(\text{NO}_3)_2$ droplet and then needle-shaped CaSO_4 crystals formed inside the droplet (Zhao et al., 2017).
151 Moreover, the amount of CaSO_4 formed in this study was much higher than that in the direct reaction of SO_2 with
152 NO_2 . The CaSO_4 solid particle constituted a significant fraction of the volume of the droplet, while in the direct
153 reaction of SO_2 with NO_2 the few needle-shaped CaSO_4 crystals that formed only constituted a small fraction of
154 the droplet volume (Zhao et al., 2017).

155 3.2 Reaction process

156 During the reaction, the amounts of carbonate, nitrate, and sulfate were determined as a function of time, as
157 shown in Fig. 3. At the beginning of the reaction, the amount of carbonate decreased slowly, while the amount of
158 nitrate and sulfate increased slowly. After a period of induction of around 50 min, the reaction accelerated
159 significantly, leading to a rapid consumption of carbonate and production of nitrate and sulfate. The decrease in
160 the amount of carbonate and the increase in the amount of nitrate was because carbonate reacted continuously
161 with NO_2 and H_2O , forming $\text{Ca}(\text{NO}_3)_2$. The detailed mechanism of the multiphase reaction of carbonate with NO_2
162 and H_2O were discussed in our previous studies (Li et al., 2010; Zhao et al., 2017). The mechanism of sulfate
163 formation is discussed in detail in Section 3.4 of the present study. Finally, the carbonate was completely
164 consumed, and the amounts of nitrate and sulfate levelled off.

165 Figure 3 shows that nitrate and sulfate were formed simultaneously during the reaction. This contrasts with
166 the observations made during the direct reaction of SO_2 with NO_2 , where nitrate was formed first, and sulfate was
167 essentially formed after the complete conversion of CaCO_3 particles to $\text{Ca}(\text{NO}_3)_2$ droplets (Zhao et al., 2017).
168 Moreover, the time taken for carbonate to be completely consumed was longer in this study than in the direct
169 reaction of SO_2 with NO_2 (~ 120 vs. ~ 40 min) when other conditions were kept the same (Zhao et al., 2017).

170 3.3 Reactive uptake coefficient of SO₂

171 The reactive uptake coefficients of SO₂ for sulfate formation (γ) in the reaction of SO₂ with the
172 O₂/NO₂/H₂O/N₂ mixture on CaCO₃ with various O₂ concentrations are shown in Table 1. The value of γ for the
173 reaction of SO₂ with O₂/NO₂ at three O₂ concentrations (5, 20, and 86%) was in the range of $(0.35\text{--}1.7) \times 10^{-5}$,
174 and was 1.2×10^{-5} in synthetic air. This latter value was 2–3 orders of magnitude higher than that for the reaction
175 of SO₂ directly with NO₂ under similar conditions (Zhao et al., 2017). When other conditions were kept constant,
176 γ increased with the O₂ concentration. This indicates that O₂ played a key role in enhancing the oxidation rate of
177 SO₂.

178 The role of O₂ in enhancing the reactive uptake of SO₂ reported here is consistent with the findings in some
179 previous studies. For example, Littlejohn et al. (1993)'s data showed that sulfite oxidation rate increases with the
180 O₂ concentration (0–5% by volume). Shen and Rochelle (1998) also found that in the presence of O₂, the aqueous
181 sulfite oxidation rate is enhanced. By investigating the oxidation of SO₂ by NO₂ in monodispersed water droplets
182 growing on carbon nuclei, Santachiara et al. (1990) found that sulfate formation rate with 2% O₂ is much higher
183 than that without O₂. Yet, our findings, as well as those in the studies referred to above, are in contrast to those
184 reported by Lee and Schwartz (1983), who found that changing from N₂ to air as a carrier gas only increases
185 bisulfite oxidation rate by no more than 10%. The difference between our study and Lee and Schwartz (1983)
186 could be due to the difference in O₂ diffusion from gas to the condensed phase and the different mechanisms
187 between the multiphase reaction on particles and the aqueous reaction.

188 Only few studies have reported the S(IV) oxidation rate in the reaction of S(IV) with O₂/NO₂ mixtures
189 (Turšič et al., 2001; Littlejohn et al., 1993). However, due to the limiting step by the aqueous phase mass transfer,
190 it is difficult to quantitatively compare the reaction rates in those studies with the uptake coefficient in our study
191 and the rate constants determined by Lee and Schwartz (1983) and Clifton et al. (1988). For example, a rate
192 constant of $2.4 \times 10^3 \text{ mol}^{-1} \text{ L s}^{-1}$ (at pH 3) can be derived from the results of Turšič et al. (2001), which is much
193 lower than the values reported by Lee and Schwartz (1983) and Clifton et al. (1988). This can be attributed to the
194 limiting step by the aqueous-phase mass transfer because the characteristic mixing time in the aqueous phase in
195 Turšič et al. (2001) was likely much longer than that of Lee and Schwartz (1983) (1.7–5.3 s), according to the
196 HSO₃⁻ concentration time series reported by Turšič et al. (2001).

197 It is important to note that the concentrations of NO₂ and SO₂ used in this study are much higher than those
198 in the ambient atmosphere. High concentrations of reactant gases are often used in laboratory studies in order to
199 simulate the ambient reactions at the time scale of days or weeks and to get high signal-to-noise ratios for
200 detecting products within minutes or hours. In the ambient atmosphere, reactive uptake coefficient of SO₂ should
201 be lower than that in this study due to the lower NO₂ concentrations when other conditions are comparable and
202 the chemical/physical processes observed in this study, such as changes in particle composition, phase,
203 hygroscopicity, and pH should be much slower due to the lower concentrations of NO₂ and SO₂.

204 3.4 Reaction mechanism

205 In the multiphase reaction of SO₂ with O₂/NO₂/H₂O on CaCO₃ particles, we found that CaCO₃ reacted with
206 NO₂ and H₂O and produced Ca(NO₃)₂, which deliquesced, forming liquid water, and provided a site for the

207 aqueous oxidation of SO₂. This process is similar to the direct reaction of SO₂ with NO₂ on CaCO₃ particles. The
208 details of this part of the reaction mechanism were discussed in our previous study (Zhao et al., 2017).

209 Once the aqueous phase was formed, SO₂ could undergo multiphase reactions with O₂/NO₂. The mechanism
210 of the direct aqueous reaction of S(IV) with NO₂ in the absence of O₂ is complex. Previous studies have proposed
211 two different mechanisms for the reaction. One involves SO₃^{•-} radical formation (Littlejohn et al., 1993; Shen and
212 Rochelle, 1998; Turšič et al., 2001) (referred as “free-radical chain” mechanism, while the other involves the
213 formation of NO₂-S(IV) complexes (Clifton et al., 1988), but no radical formation (referred as “NO₂-S(IV)
214 complex” mechanism).

215 According to the NO₂-S(IV) ~~adduct-complex~~ mechanism, the presence of O₂ should not affect the SO₂
216 oxidation rate; however, in this study, a substantial enhancement in the SO₂ oxidation rate was observed in the
217 presence of O₂ compared with that in the absence of O₂. Therefore, the NO₂-S(IV) ~~adduct-complex~~ mechanism
218 was ~~not considered less likely~~ to have been important in this study.

219 In the free-radical chain mechanism, the SO₃^{•-} radical is proposed to be formed (~~R7R8~~, Table 2), which is
220 based on the observation of S₂O₆²⁻ formation, with S₂O₆²⁻ known to be the combination reaction product of SO₃^{•-}
221 (Eriksen, 1974; Hayon et al., 1972; Deister and Warneck, 1990; Brandt et al., 1994; Waygood and McElroy,
222 1992). In addition to SO₄²⁻ and NO₂⁻, S₂O₆²⁻ was detected with an appreciable yield using Raman spectroscopy,
223 following the reaction of NO₂ with aqueous sulfite (Littlejohn et al., 1993). S₂O₆²⁻ was also observed in the
224 aqueous oxidation of bisulfite in an N₂-saturated solution in the presence of Fe(III) using ion-interaction
225 chromatography (Podkrajšek et al., 2002). The SO₃^{•-} radical can react via two pathways, forming either S₂O₆²⁻ or
226 SO₄²⁻ (~~R8R9-R10R11~~, Table 2). The reactions ~~R8R9-R10-R11~~ have been well established in studies of S(IV)
227 oxidation by other pathways, including OH oxidation, photo-oxidation, and transition metal catalyzed oxidation
228 (Eriksen, 1974; Hayon et al., 1972; Deister and Warneck, 1990; Brandt et al., 1994; Brandt and Vaneldik, 1995;
229 Waygood and McElroy, 1992). In addition, although previous studies have not reported the direct observation of
230 the SO₃^{•-} radical in the aqueous reaction of S(IV) with NO₂, SO₃^{•-} was observed in the reaction of NO₂⁻ with
231 SO₃²⁻ in an acidic buffer solution (pH = 4.0) using electron spin resonance (ESR) (Shi, 1994). Because NO₂⁻ is
232 formed in the aqueous reaction of SO₂ with NO₂, and S₂O₆²⁻ as the combination reaction product of SO₃^{•-} is
233 observed (Littlejohn et al., 1993), SO₃^{•-} formation is plausible.

234 In the presence of O₂, the SO₃^{•-} radical can react rapidly with O₂, forming the SO₅^{•-} radical (~~R11R12~~, Table
235 2). Following this reaction, a number of chain reactions can occur to ultimately form sulfate (Littlejohn et al.,
236 1993; Seinfeld and Pandis, 2006; Shen and Rochelle, 1998) (~~R12R13-R15R16~~, Table 2). Littlejohn et al. (1993)
237 observed that the amount of S₂O₆²⁻ relative to SO₄²⁻ formed in the aqueous reaction of NO₂ with sulfite decreases
238 in the presence of O₂ compared with the reaction in the absence of O₂. At low NO₂ concentrations (< 5 ppm),
239 S₂O₆²⁻ is undetectable in the presence of O₂. This indicates that O₂ suppresses the reaction pathway of S₂O₆²⁻
240 formation (~~R8R9~~, Table 2). Because the SO₃^{•-} radical can react rapidly with O₂, forming the SO₅^{•-} radical, and
241 would therefore be consumed, the suppression of S₂O₆²⁻ formation can be attributed to the reaction of SO₃^{•-} with
242 O₂ (~~R11R12~~, Table 2). The reactions ~~R11R12-R15-R16~~ have been well established by studies of the oxidation of
243 S(IV) by OH or photo-oxidation, and all the radicals have been observed (Hayon et al., 1972; Huie et al., 1989;
244 Huie and Neta, 1987; Chameides and Davis, 1982; Seinfeld and Pandis, 2006).

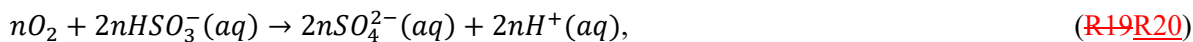
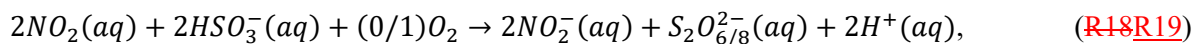
245 The free-radical chain mechanism is consistent with the findings of this study and is therefore more plausible.
246 The enhancement of the SO₂ oxidation rate in the reaction of SO₂ with O₂/NO₂/H₂O on CaCO₃ particles
247 compared with that in the direct reaction of SO₂ with NO₂/H₂O was attributed to O₂. Although during the reaction
248 in the absence of O₂—i.e., the direct oxidation of SO₂ by NO₂—the SO₃^{•-} radical can be formed (R7), the reaction
249 chain cannot propagate (R11R12–R15R16). Therefore, the S(IV) oxidation rate and the reactive uptake
250 coefficient of SO₂ were much lower than that in the presence of O₂. According to the difference between the
251 reactive uptake coefficient in this study and in the direct reaction of SO₂ with NO₂ (Zhao et al., 2017), the sulfate
252 production rate via chain reactions due to the presence of O₂ (20%) was 2–3 orders of magnitude faster than the
253 direct oxidation of SO₂ by NO₂. This indicates that sulfate production in the reaction of SO₂ with O₂/NO₂ was
254 largely due to O₂ oxidation via the chain reaction pathway, i.e., “autoxidation” of S(IV), rather than the direct
255 oxidation of SO₂ by NO₂ and thus O₂ was the main oxidant of SO₂.

256 In addition to the two mechanisms above, Spindler et al. (2003) proposed a reaction mechanism involving
257 first NO₂-S(IV) complex formation and subsequent SO₃^{•-} radical formation (R3, R7). NO₂-S(IV) complex may
258 establish an equilibrium with SO₃^{•-} in contrast to the direct formation of SO₃^{•-} via the reaction of NO₂ with SO₂.
259 Higher concentration of O₂ favors the conversion of SO₃^{•-} to SO₅^{•-} and thus can reduce the SO₃^{•-} concentration,
260 shifting the equilibrium to the product side and promoting the overall S(IV) oxidation. O₂ can act in a similar way
261 as in the free-radical chain mechanism. Admittedly, we cannot rule out the possibility NO₂-S(IV) complex
262 formation. But such a mechanism can still be classified as the free-radical chain mechanism since the S(IV)
263 oxidation still proceeds via the radical chain reactions.

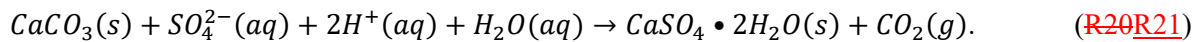
264 Although the direct oxidation of SO₂ by NO₂ only accounted for a very small fraction of sulfate formation,
265 NO₂ played an important role in the SO₂ oxidation by initiating the chain reactions via the production of the SO₃^{•-}
266 radical (R7). In the experiment without NO₂, but with other reaction conditions the same, we were unable to
267 detect sulfate after 5 h of reaction. This indicates that O₂ by itself cannot initiate the chain reactions (although it
268 favors chain propagation), and that the oxidation of SO₂ by O₂ was slow. The effect on the SO₂ oxidation rate
269 when both NO₂ and O₂ were present was much higher than the sum of the effect of NO₂ and O₂. We refer to this
270 effect as the synergy of NO₂ and O₂, which resulted in the fast oxidation of SO₂ to form sulfate in this study. This
271 effect is similar to a “ternary” reaction found with the reaction of NO₂-particles-H₂O or SO₂-particles-O₃ (Zhu et
272 al., 2011), where the reaction rate can be much faster than the sum of the reaction rates for the reaction of the
273 second and third reactant with the first reactant. In addition to acting as the initiator of chain reactions, NO₂ also
274 contributed to the formation of the aqueous phase through the reaction with CaCO₃, forming Ca(NO₃)₂ as
275 discussed above, which provided a site for S(IV) oxidation.

276 Based on the discussion above, we summarize the reaction mechanism that occurred in this study in Table 2.
277 The reactions are classified as chain initiation, chain propagation, and chain termination. The dominant S(IV)
278 species depends on pH. Due to the fast dissociations of SO₂•H₂O and HSO₃⁻, reactions consuming one of these
279 S(IV) species will result in instantaneous re-establishment of the equilibria between them (Seinfeld and Pandis,
280 2006). In this study, the pH of the aqueous layer of Ca(NO₃)₂ may change dynamically with time during the
281 reaction and may not be completely homogeneous within the aqueous droplet. The pH values could vary between
282 ~3 and ~7.6. In the surface of the aqueous layer, pH was mainly determined by the gas-aqueous equilibrium of
283 SO₂, and was estimated to be ~3. In the vicinity of the CaCO₃ core, pH was mainly determined by the hydrolysis

284 of carbonate and was estimated to be ~7.6. It is likely that both HSO_3^- and SO_3^{2-} were present, and the dominant
 285 species depended on the reaction time and location within the aqueous droplet. Nevertheless, to make the reaction
 286 mechanism clearer, HSO_3^- was used in the reaction equations. Similar reaction equations are also applicable to
 287 SO_3^{2-} because of the fast dissociations of $\text{SO}_2 \cdot \text{H}_2\text{O}$ and HSO_3^- . Overall, the reaction can be written as follows,
 288 which clearly shows that O_2 was the main oxidant for sulfate formation:



291 where $n \gg 1$. Once sulfuric acid was formed, it reacted with CaCO_3 , forming CaSO_4 :



293 As mentioned above, compared with the direct reaction of SO_2 with NO_2 , CaCO_3 was consumed more
 294 slowly in the reaction with O_2/NO_2 . There were two possible reasons for this. First, the $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$ formed in
 295 the reaction could cover the CaCO_3 surface and partly suppress the diffusion of aqueous ions, such as protons,
 296 and also limit the contact of reactants with the surface of the CaCO_3 particles, thus reducing the CaCO_3
 297 consumption rate. Second, compared with the direct reaction of SO_2 with NO_2 , a much higher fraction of CaCO_3
 298 was converted to $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$ instead of $\text{Ca}(\text{NO}_3)_2$ due to the fast production of $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$. Therefore, the
 299 volume of a $\text{Ca}(\text{NO}_3)_2$ droplet was much smaller than that in the direct reaction of SO_2 with NO_2 for a given
 300 CaCO_3 particle. Because the uptake rate of NO_2 was proportional to the droplet surface area and the NO_2
 301 hydrolysis rate was proportional to the droplet volume, the rate of nitric acid production from NO_2 hydrolysis and
 302 its reaction rate with CaCO_3 were reduced. Therefore, the CaCO_3 particles were consumed more slowly in the
 303 reaction with O_2/NO_2 .

304 4 Conclusion and implications

305 We investigated the multiphase reaction of SO_2 with $\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$ on CaCO_3 particles. The reaction
 306 converted CaCO_3 particles to $\text{Ca}(\text{NO}_3)_2$ droplets, in which $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$ was embedded and accounted for a
 307 significant fraction of the droplet volume by the end of the reaction. The $\text{Ca}(\text{NO}_3)_2$ droplet formed by the reaction
 308 of CaCO_3 with NO_2 provided a site for the multiphase oxidation of SO_2 . Generally, nitrate and sulfate were
 309 formed simultaneously. The reactive uptake coefficient of SO_2 for sulfate formation in the reaction of SO_2 with
 310 $\text{NO}_2/\text{H}_2\text{O}$ in synthetic air was determined to be around 10^{-5} . Compared with the reaction of SO_2 with NO_2 on a
 311 CaCO_3 particle in the absence of O_2 , i.e., the direct oxidation of SO_2 by NO_2 in N_2 , sulfate production rate in the
 312 reaction of SO_2 with O_2/NO_2 was enhanced by 2–3 orders of magnitude. According to the findings of this study
 313 and the existing literature, SO_2 oxidation likely proceeded via a free-radical chain reaction mechanism. O_2 was
 314 the main oxidant of SO_2 , and NO_2 mainly acted as an initiator of the chain reactions. The synergy of NO_2 and O_2
 315 resulted in the fast oxidation of SO_2 . The absence of either NO_2 or O_2 led to much slower SO_2 oxidation.

316 Using a method developed in our previous study (Zhao et al., 2017), we assessed the importance of the
 317 multiphase oxidation of SO_2 by O_2 in the presence of NO_2 by estimating the lifetime of SO_2 due to multiphase
 318 reactions and the lifetime due to the gas phase reaction (with the OH radical). The lifetime of SO_2 due to the
 319 multiphase reaction of SO_2 with O_2/NO_2 was estimated to be around 20 days using the reactive uptake coefficient
 320 of SO_2 (1.2×10^{-5}) and the typical particle surface area concentration for mineral aerosols in winter in Beijing

321 $(6.3 \times 10^{-6} \text{ cm}^2 \text{ cm}^{-3})$ (Huang et al., 2015). This lifetime is comparable to the lifetime of SO_2 due to the gas phase
322 reaction with OH, which is ~ 12 days assuming that the daytime OH concentration is 1×10^6 molecules cm^{-3}
323 (Lelieveld et al., 2016; Prinn et al., 2005). Therefore, we conclude that the multiphase oxidation of SO_2 by O_2 in
324 the presence of NO_2 is likely to be an important source of sulfate and a sink of SO_2 in the ambient atmosphere,
325 and can play a significant role in the sulfate formation during severe haze episodes, such as those that frequently
326 occur in China. During haze episodes, there are high concentrations of SO_2 and NO_2 and relative humidity is
327 often high (Zhang et al., 2014; Wang et al., 2016; Zheng et al., 2015b). Under these conditions, the multiphase
328 oxidation of SO_2 by O_2 in the presence of NO_2 could proceed rapidly, forming sulfate. The enhanced sulfate
329 concentration due to multiphase reactions and resulting aerosol water content can further promote the multiphase
330 oxidation of SO_2 . The reaction thus proceeds in a self-accelerating way. Therefore, it can contribute significantly
331 to sulfate formation during haze episodes, which could explain the discrepancies between the observed and
332 modelled sulfate concentrations (Cheng et al., 2016; Gao et al., 2016; Wang et al., 2016; Zheng et al., 2015a).

333 In addition, elucidating the mechanism of the multiphase reaction of SO_2 with $\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$ in the
334 atmosphere is important for the other atmospheric implications of the reaction besides sulfate formation.
335 According to the reaction mechanism, the direct oxidation of SO_2 by NO_2 forms sulfate and nitrite, with a
336 stoichiometry of 1:1, and nitrite can further form HONO under acidic conditions. The HONO could then
337 evaporate into the atmosphere, where it would be an important source of OH radical. If NO_2 were the main
338 oxidant of SO_2 in the multiphase reaction, the reaction would form one HONO molecule for every sulfate formed.
339 Thus, the oxidation of SO_2 by NO_2 can simultaneously be an important source of HONO and OH radical, and
340 SO_2 oxidation would be strongly coupled with reactive nitrogen chemistry. However, according to the
341 mechanism of this study, NO_2 only acted as an initiator of the chain reactions in SO_2 oxidation and essentially all
342 the SO_2 was oxidized by O_2 . Therefore, the amount of HONO formation per sulfate formed was trivial. The
343 oxidation of SO_2 by O_2/NO_2 is expected to be neither an important source of HONO and OH in the atmosphere
344 nor to have a significant influence on reactive nitrogen chemistry.

345 In this study, we investigated the reaction of SO_2 with O_2 in the presence of NO_2 at three O_2 concentrations.
346 The influence of the O_2 concentration was shown to be significant. Future experimental results with more O_2
347 concentration levels would provide more insights into the reaction mechanism and process.

348 In addition, in the ambient atmosphere, the internal mixing of organics with S(IV) in particles may influence
349 the S(IV) oxidation rate by O_2 in the presence of NO_2 . When organics is abundant in particles, for example during
350 haze episodes in China, it can react with and thus scavenge radical anion carriers such as $\text{SO}_5^{\cdot-}$ and $\text{SO}_4^{\cdot-}$
351 (Herrmann, 2003; Herrmann et al., 2015; Huie, 1995). Therefore, the presence of internally mixed organics can
352 reduce the effectivity of the potential radical reaction chain and of S(IV) oxidation, which can undermine the
353 importance of the oxidation by O_2 in the presence of NO_2 in the overall S(IV) oxidation.

354 Acknowledgements

355 This work was supported by Natural Science Foundation Committee of China (41421064, 21190051,
356 40490265, 91544000) and Ministry of Science and Technology (Grant No. 2002CB410802).

357 **References**

- 358 Brandt, C., Fabian, I., and Vaneldik, R.: Kinetics and mechanism of the iron(III)-catalyzed autoxidation of
359 sulfur(IV) oxides in aqueous-solution - evidence for the redox cycling of iron in the presence of oxygen and
360 modeling of the overall reaction-mechanism, *Inorg. Chem.*, 33, 687-701, 10.1021/ic00082a012, 1994.
- 361 Brandt, C., and Vaneldik, R.: Transition metal-catalyzed oxidation of sulfur (IV) oxides. Atmospheric-relevant
362 processes and mechanisms, *Chem. Rev.*, 95, 119-190, 10.1021/cr00033a006, 1995.
- 363 Chameides, W. L., and Davis, D. D.: The free-radical chemistry of cloud droplets and its impact upon the
364 composition of rain, *J. Geophys. Res.-Oceans*, 87, 4863-4877, 10.1029/JC087iC07p04863, 1982.
- 365 Cheng, Y. F., Zheng, G. J., Wei, C., Mu, Q., Zheng, B., Wang, Z. B., Gao, M., Zhang, Q., He, K. B., Carmichael,
366 G., Poschl, U., and Su, H.: Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events
367 in China, *Sci. Adv.*, 2, 10.1126/sciadv.1601530, 2016.
- 368 Clifton, C. L., Altstein, N., and Huie, R. E.: Rate-constant for the reaction of NO₂ with sulfur(IV) over the pH
369 range 5.3-13, *Environ. Sci. Technol.*, 22, 586-589, 10.1021/es00170a018, 1988.
- 370 Davidovits, P., Kolb, C. E., Williams, L. R., Jayne, J. T., and Worsnop, D. R.: Mass accommodation and
371 chemical reactions at gas-liquid interfaces, *Chem. Rev.*, 106, 1323-1354, 10.1021/cr040366k, 2006.
- 372 Deister, U., and Warneck, P.: Photooxidation of sulfite (SO₃²⁻) in aqueous solution, *J. Phys. Chem.*, 94,
373 2191-2198, 10.1021/j100368a084, 1990.
- 374 Eriksen, T. E.: pH Effects on the pulse radiolysis of deoxygenated aqueous solutions of sulphur dioxide, *Journal*
375 *of the Chemical Society, Faraday Transactions 1: Physical Chemistry in Condensed Phases*, 70, 208-215,
376 10.1039/f19747000208, 1974.
- 377 Gao, M., Carmichael, G. R., Wang, Y., Ji, D., Liu, Z., and Wang, Z.: Improving simulations of sulfate aerosols
378 during winter haze over Northern China: the impacts of heterogeneous oxidation by NO₂, *Front. Environ. Sci.*
379 *Eng.*, 10, 16, 10.1007/s11783-016-0878-2, 2016.
- 380 Guo, H., Weber, R. J., and Nenes, A.: High levels of ammonia do not raise fine particle pH sufficiently to yield
381 nitrogen oxide-dominated sulfate production, *Sci. Rep.*, 7, 12109, 10.1038/s41598-017-11704-0, 2017.
- 382 Hayon, E., Treinin, A., and Wilf, J.: Electronic spectra, photochemistry, and autoxidation mechanism of the
383 sulfite-bisulfite-pyrosulfite systems. SO₂⁻, SO₃⁻, SO₄⁻, and SO₅⁻ radicals, *J. Am. Chem. Soc.*, 94, 47-57,
384 10.1021/ja00756a009, 1972.
- 385 He, H., Wang, Y., Ma, Q., Ma, J., Chu, B., Ji, D., Tang, G., Liu, C., Zhang, H., and Hao, J.: Mineral dust and
386 NO_x promote the conversion of SO₂ to sulfate in heavy pollution days, *Sci. Rep.*, 4, 10.1038/srep04172, 2014.
- 387 Herrmann, H.: Kinetics of aqueous phase reactions relevant for atmospheric chemistry, *Chem. Rev.*, 103,
388 4691-4716, 10.1021/cr020658q, 2003.
- 389 Herrmann, H., Schaefer, T., Tilgner, A., Styler, S. A., Weller, C., Teich, M., and Otto, T.: Tropospheric
390 Aqueous-Phase Chemistry: Kinetics, Mechanisms, and Its Coupling to a Changing Gas Phase, *Chem. Rev.*, 115,
391 4259-4334, 10.1021/cr500447k, 2015.
- 392 Huang, L., Zhao, Y., Li, H., and Chen, Z.: Kinetics of Heterogeneous Reaction of Sulfur Dioxide on Authentic
393 Mineral Dust: Effects of Relative Humidity and Hydrogen Peroxide, *Environ. Sci. Technol.*, 49, 10797-10805,
394 10.1021/acs.est.5b03930, 2015.

395 Huie, R. E., and Neta, P.: Kinetics of one-electron transfer-reactions involving ClO₂ and NO₂, *J. Phys. Chem.*,
396 90, 1193-1198, 10.1021/j100278a046, 1986.

397 Huie, R. E., and Neta, P.: Rate constants for some oxidations of S(IV) by radicals in aqueous-solutions, *Atmos.*
398 *Environ.*, 21, 1743-1747, 10.1016/0004-6981(87)90113-2, 1987.

399 Huie, R. E., Clifton, C. L., and Altstein, N.: A pulse radiolysis and flash photolysis study of the radicals SO₂,
400 SO₃, SO₄ and SO₅, *Radiat. Phys. Chem.*, 33, 361-370, 1989.

401 Huie, R. E.: Free radical chemistry of the atmospheric aqueous phase, in: *Progress and Problems in Atmospheric*
402 *Chemistry*, WORLD SCIENTIFIC, 374-419, 1995.

403 Lee, Y.-N., and Schwartz, S. E.: Kinetics of oxidation of aqueous sulfur (IV) by nitrogen dioxide, in:
404 *Precipitation Scavenging, Dry Deposition and Resuspension*, edited by: Pruppacher, H. R., Semonin, R. G., and
405 Slinn, W. G. N., Elsevier, New York, 453-466, 1983.

406 Lelieveld, J., Gromov, S., Pozzer, A., and Taraborrelli, D.: Global tropospheric hydroxyl distribution, budget and
407 reactivity, *Atmos. Chem. Phys.*, 16, 12477-12493, 10.5194/acp-16-12477-2016, 2016.

408 Li, H. J., Zhu, T., Zhao, D. F., Zhang, Z. F., and Chen, Z. M.: Kinetics and mechanisms of heterogeneous
409 reaction of NO₂ on CaCO₃ surfaces under dry and wet conditions, *Atmos. Chem. Phys.*, 10, 463-474, 2010.

410 Littlejohn, D., Wang, Y. Z., and Chang, S. G.: Oxidation of aqueous sulfite ion by nitrogen-dioxide, *Environ. Sci.*
411 *Technol.*, 27, 2162-2167, 10.1021/es00047a024, 1993.

412 Liu, M. X., Song, Y., Zhou, T., Xu, Z. Y., Yan, C. Q., Zheng, M., Wu, Z. J., Hu, M., Wu, Y. S., and Zhu, T.: Fine
413 particle pH during severe haze episodes in northern China, *Geophys. Res. Lett.*, 44, 5213-5221,
414 10.1002/2017gl073210, 2017.

415 Liu, Y. J., Zhu, T., Zhao, D. F., and Zhang, Z. F.: Investigation of the hygroscopic properties of Ca(NO₃)(2) and
416 internally mixed Ca(NO₃)(2)/CaCO₃ particles by micro-Raman spectrometry, *Atmos. Chem. Phys.*, 8,
417 7205-7215, 2008.

418 Ma, Q., He, H., Liu, Y., Liu, C., and Grassian, V. H.: Heterogeneous and multiphase formation pathways of
419 gypsum in the atmosphere, *Phys. Chem. Chem. Phys.*, 15, 19196-19204, 10.1039/c3cp53424c, 2013.

420 Nakamoto, K.: *Infrared and Raman Spectra of Inorganic and Coordination Compounds Part A*, John Wiley
421 & Sons, New York, 221-247 pp., 1997.

422 Nash, T.: Effect of nitrogen-dioxide and of some transition-metals on the oxidation of dilute bisulfite solutions,
423 *Atmos. Environ.*, 13, 1149-1154, 10.1016/0004-6981(79)90038-6, 1979.

424 Pöschl, U., Rudich, Y., and Ammann, M.: Kinetic model framework for aerosol and cloud surface chemistry and
425 gas-particle interactions - Part 1: General equations, parameters, and terminology, *Atmos. Chem. Phys.*, 7,
426 5989-6023, 2007.

427 Podkrajšek, B., Grgić, I., and Turšič, J.: Determination of sulfur oxides formed during the S(IV) oxidation in the
428 presence of iron, *Chemosphere*, 49, 271-277, [https://doi.org/10.1016/S0045-6535\(02\)00324-7](https://doi.org/10.1016/S0045-6535(02)00324-7), 2002.

429 Prinn, R. G., Huang, J., Weiss, R. F., Cunnold, D. M., Fraser, P. J., Simmonds, P. G., McCulloch, A., Harth, C.,
430 Reimann, S., Salameh, P., O'Doherty, S., Wang, R. H. J., Porter, L. W., Miller, B. R., and Krummel, P. B.:
431 Evidence for variability of atmospheric hydroxyl radicals over the past quarter century, *Geophys. Res. Lett.*, 32,
432 10.1029/2004gl022228, 2005.

433 Santachiara, G., Prodi, F., and Vivarelli, F.: SO₂ oxidation in monodisperse droplets grown on carbon nuclei in
434 presence of NO₂, *J. Aerosol Sci.*, 21, S221-S224, 10.1016/0021-8502(90)90224-1, 1990.

435 Santachiara, G., Prodi, F., and Vivarelli, F.: Further experiments on SO₂ oxidation rate in monodisperse droplets
436 grown on carbon nuclei in presence of O₂ and NO₂, *J. Aerosol Sci.*, 24, 683-685,
437 10.1016/0021-8502(93)90024-4, 1993.

438 Sarma, L. P., Prasad, P. S. R., and Ravikumar, N.: Raman spectroscopic study of phase transitions in natural
439 gypsum, *J. Raman Spectrosc.*, 29, 851-856, 10.1002/(sici)1097-4555(199809)29:9<851::aid-jrs313>3.0.co;2-s,
440 1998.

441 Seinfeld, J. H., and Pandis, S. N.: *Atmospheric chemistry and physics: from air pollution to climate change*, 2nd
442 ed., John Wiley & Sons. Inc., 2006.

443 Shen, C. H., and Rochelle, G. T.: Nitrogen Dioxide Absorption and Sulfite Oxidation in Aqueous Sulfite, *Environ.*
444 *Sci. Technol.*, 32, 1994-2003, 10.1021/es970466q, 1998.

445 Shi, X. L.: Generation of SO₃⁻ and OH radicals in SO₃²⁻ reactions with inorganic environmental-pollutants and its
446 implications to SO₃²⁻ toxicity, *J. Inorg. Biochem.*, 56, 155-165, 10.1016/0162-0134(94)85002-x, 1994.

447 Spindler, G., Hesper, J., Brüggemann, E., Dubois, R., Müller, T., and Herrmann, H.: Wet annular denuder
448 measurements of nitrous acid: laboratory study of the artefact reaction of NO₂ with S(IV) in aqueous solution
449 and comparison with field measurements, *Atmos. Environ.*, 37, 2643-2662, 10.1016/s1352-2310(03)00209-7,
450 2003.

451 Turšič, J., Grgić, I., and Bizjak, M.: Influence of NO₂ and dissolved iron on the S(IV) oxidation in synthetic
452 aqueous solution, *Atmos. Environ.*, 35, 97-104, [https://doi.org/10.1016/S1352-2310\(00\)00283-1](https://doi.org/10.1016/S1352-2310(00)00283-1), 2001.

453 Wang, G., Zhang, R., Gomez, M. E., Yang, L., Levy Zamora, M., Hu, M., Lin, Y., Peng, J., Guo, S., Meng, J., Li,
454 J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J., Cao, J., An, Z., Zhou, W., Li, G., Wang, J., Tian, P.,
455 Marrero-Ortiz, W., Secrest, J., Du, Z., Zheng, J., Shang, D., Zeng, L., Shao, M., Wang, W., Huang, Y., Wang, Y.,
456 Zhu, Y., Li, Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F., Rosenfeld, D., Liss, P. S., Duce, R. A., Kolb,
457 C. E., and Molina, M. J.: Persistent sulfate formation from London Fog to Chinese haze, *Proc. Nat. Acad. Sci.*
458 *U.S.A.*, 113, 13630-13635, 10.1073/pnas.1616540113, 2016.

459 Waygood, S. J., and McElroy, W. J.: Spectroscopy and decay kinetics of the sulfite radical anion in aqueous
460 solution, *J. Chem. Soc.-Faraday Trans.*, 88, 1525-1530, 10.1039/ft9928801525, 1992.

461 Xue, J., Yuan, Z. B., Griffith, S. M., Yu, X., Lau, A. K. H., and Yu, J. Z.: Sulfate Formation Enhanced by a
462 Cocktail of High NO_x, SO₂, Particulate Matter, and Droplet pH during Haze-Fog Events in Megacities in China:
463 An Observation-Based Modeling Investigation, *Environ. Sci. Technol.*, 50, 7325-7334, 10.1021/acs.est.6b00768,
464 2016.

465 Zhang, J. K., Sun, Y., Liu, Z. R., Ji, D. S., Hu, B., Liu, Q., and Wang, Y. S.: Characterization of submicron
466 aerosols during a month of serious pollution in Beijing, 2013, *Atmos. Chem. Phys.*, 14, 2887-2903,
467 10.5194/acp-14-2887-2014, 2014.

468 Zhao, D., Song, X., Zhu, T., Zhang, Z., and Liu, Y.: Multiphase Reaction of SO₂ with NO₂ on CaCO₃ Particles.
469 1. Oxidation of SO₂ by NO₂, *Atmos. Chem. Phys. Discuss.*, 2017, 1-23, 10.5194/acp-2017-610, 2017.

470 Zhao, D. F., Zhu, T., Chen, Q., Liu, Y. J., and Zhang, Z. F.: Raman micro-spectrometry as a technique for
471 investigating heterogeneous reactions on individual atmospheric particles, *Sci. China Chem.*, 54, 154-160,
472 10.1007/s11426-010-4182-x, 2011.

473 Zheng, B., Zhang, Q., Zhang, Y., He, K. B., Wang, K., Zheng, G. J., Duan, F. K., Ma, Y. L., and Kimoto, T.:
474 Heterogeneous chemistry: a mechanism missing in current models to explain secondary inorganic aerosol
475 formation during the January 2013 haze episode in North China, *Atmos. Chem. Phys.*, 15, 2031-2049,
476 10.5194/acp-15-2031-2015, 2015a.

477 Zheng, G. J., Duan, F. K., Su, H., Ma, Y. L., Cheng, Y., Zheng, B., Zhang, Q., Huang, T., Kimoto, T., Chang, D.,
478 Poschl, U., Cheng, Y. F., and He, K. B.: Exploring the severe winter haze in Beijing: the impact of synoptic
479 weather, regional transport and heterogeneous reactions, *Atmos. Chem. Phys.*, 15, 2969-2983,
480 10.5194/acp-15-2969-2015, 2015b.

481 Zhu, T., Shang, J., and Zhao, D. F.: The roles of heterogeneous chemical processes in the formation of an air
482 pollution complex and gray haze, *Sci. China Chem.*, 54, 145-153, 10.1007/s11426-010-4181-y, 2011.

483

484

485

486

Table 1. Reactive uptake coefficient of SO₂ for sulfate formation at 82% RH and at different O₂ concentrations.

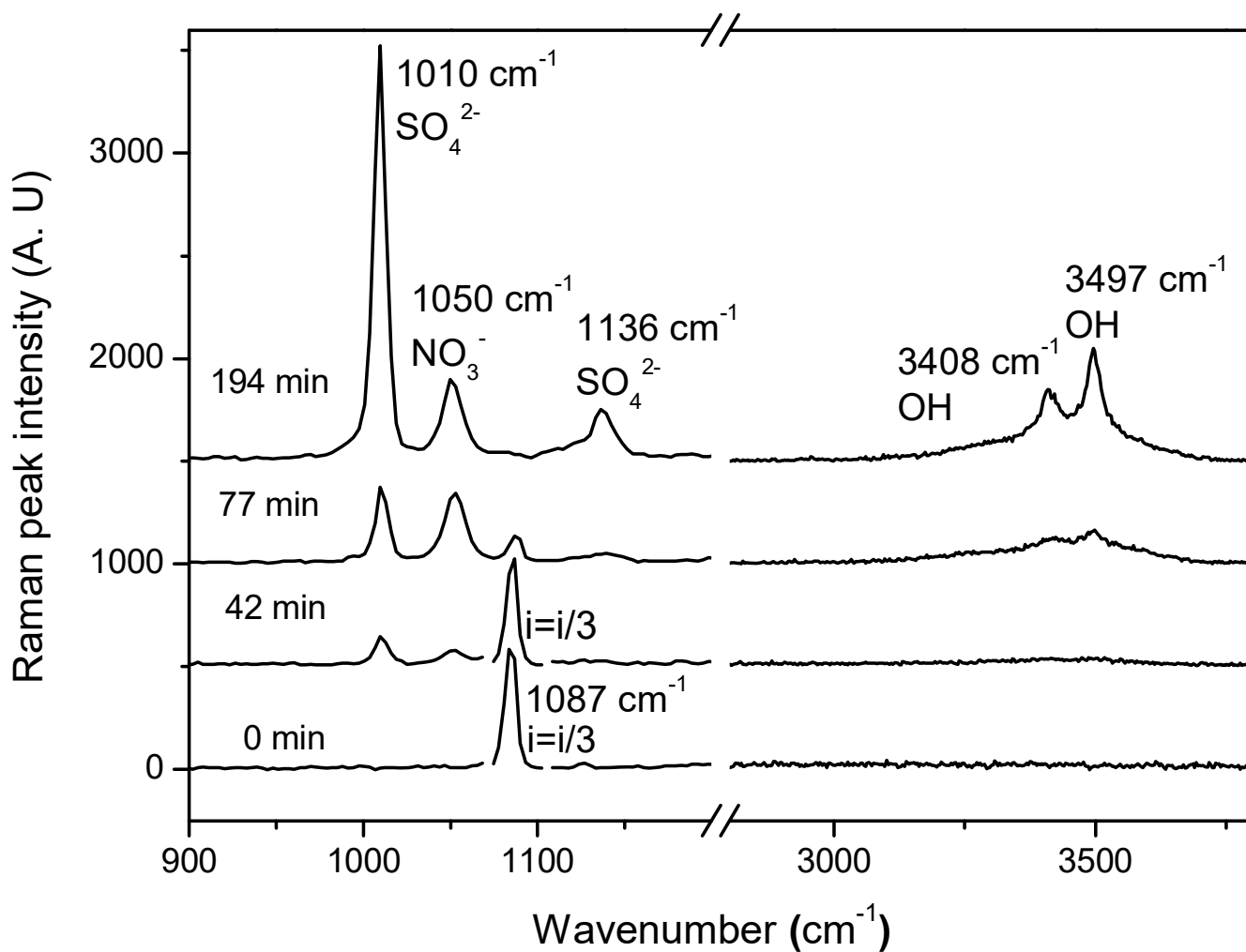
SO ₂ /NO ₂ /O ₂ concentration	γ
75 ppm/ 75 ppm/ 86 %	1.7×10^{-5}
75 ppm/ 75 ppm/ 20 %	1.2×10^{-5}
75 ppm/ 75 ppm/ 5 %	3.5×10^{-6}

487
488

489

Table 2. Summary of the mechanism of the reaction S(IV) with O₂/NO₂

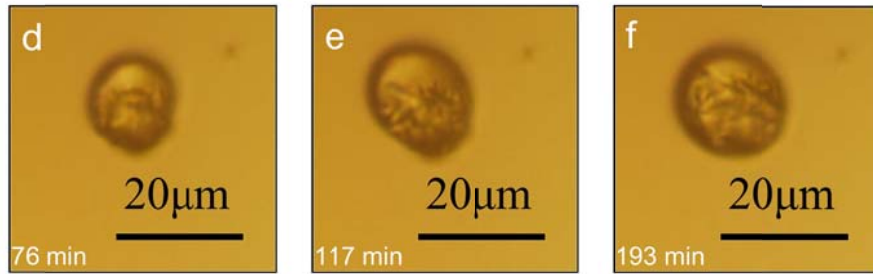
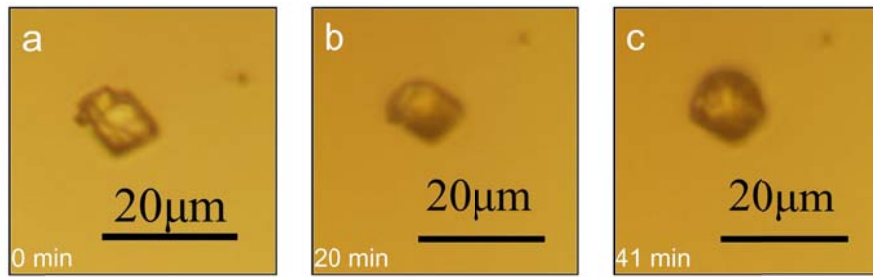
Step	Reactions
Initiation	$\text{NO}_2(\text{aq}) + \text{HSO}_3^-(\text{aq}) \rightarrow \text{NO}_2^-(\text{aq}) + \text{SO}_3^{\bullet-}(\text{aq}) + \text{H}^+(\text{aq})$ (R7R8)
Propagation	$\text{SO}_3^{\bullet-}(\text{aq}) + \text{O}_2(\text{aq}) \rightarrow \text{SO}_5^{\bullet-}(\text{aq})$ (R11R12)
	$\text{SO}_5^{\bullet-}(\text{aq}) + \text{HSO}_3^-(\text{aq}) \rightarrow \text{HSO}_5^-(\text{aq}) + \text{SO}_3^{\bullet-}(\text{aq})$ (R12R13)
	$\text{HSO}_5^-(\text{aq}) + \text{HSO}_3^-(\text{aq}) \rightarrow 2\text{SO}_4^{2-}(\text{aq}) + 2\text{H}^+(\text{aq})$ (R13R14)
	$\text{SO}_5^{\bullet-}(\text{aq}) + \text{HSO}_3^-(\text{aq}) \rightarrow \text{SO}_4^{2-}(\text{aq}) + \text{SO}_4^{\bullet-}(\text{aq}) + \text{H}^+(\text{aq})$ (R14R15)
	$\text{SO}_4^{\bullet-}(\text{aq}) + \text{HSO}_3^-(\text{aq}) \rightarrow \text{SO}_4^{2-}(\text{aq}) + \text{SO}_3^{\bullet-}(\text{aq}) + \text{H}^+(\text{aq})$ (R15R16)
Termination	$\text{SO}_3^{\bullet-}(\text{aq}) + \text{SO}_3^{\bullet-}(\text{aq}) \rightarrow \text{S}_2\text{O}_6^{2-}(\text{aq})$ (R8R9)
	$\text{SO}_3^{\bullet-}(\text{aq}) + \text{SO}_3^{\bullet-}(\text{aq}) \rightarrow \text{SO}_3^{2-}(\text{aq}) + \text{SO}_3$ (R109)
	$\text{SO}_3(\text{aq}) + \text{H}_2\text{O} \rightarrow \text{SO}_4^{2-}(\text{aq}) + 2\text{H}^+(\text{aq})$ (R10R11)
	$\text{SO}_4^{\bullet-}(\text{aq}) + \text{SO}_4^{\bullet-}(\text{aq}) \rightarrow \text{S}_2\text{O}_8^{2-}(\text{aq})$ (R16R17)
	$\text{SO}_5^{\bullet-}(\text{aq}) + \text{SO}_5^{\bullet-}(\text{aq}) \rightarrow \text{S}_2\text{O}_8^{2-}(\text{aq}) + \text{O}_2(\text{aq})$ (R17R18)



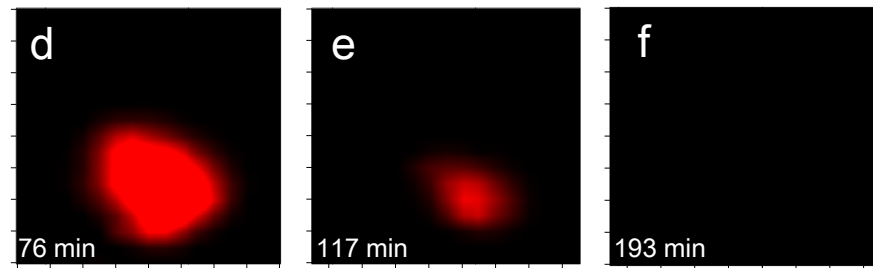
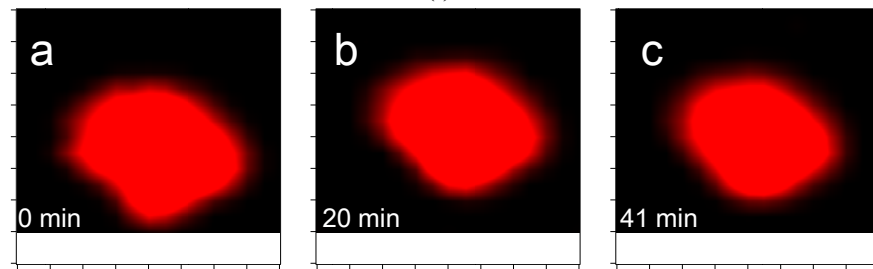
492

493 Figure 1. Raman spectra of a CaCO_3 particle during the multiphase reaction of SO_2 with $\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$
 494 on the particle. SO_2 : 75 ppm, NO_2 : 75 ppm, RH: 72%, O_2 : 20%. The peak intensity of carbonate (1087
 495 cm^{-1}) at 0 and 42 min was divided by three for clarity.

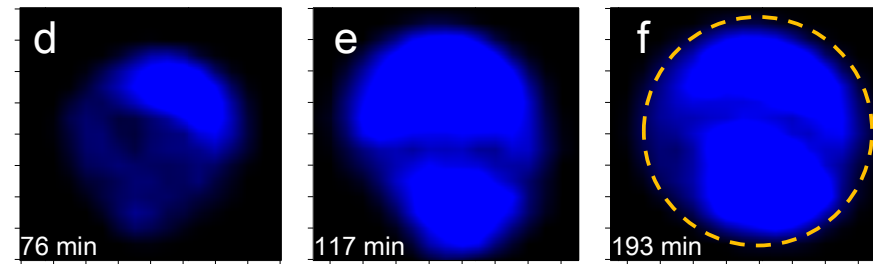
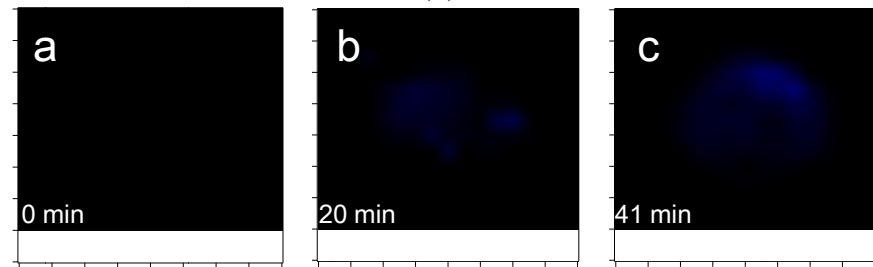
496



(i)



(ii)

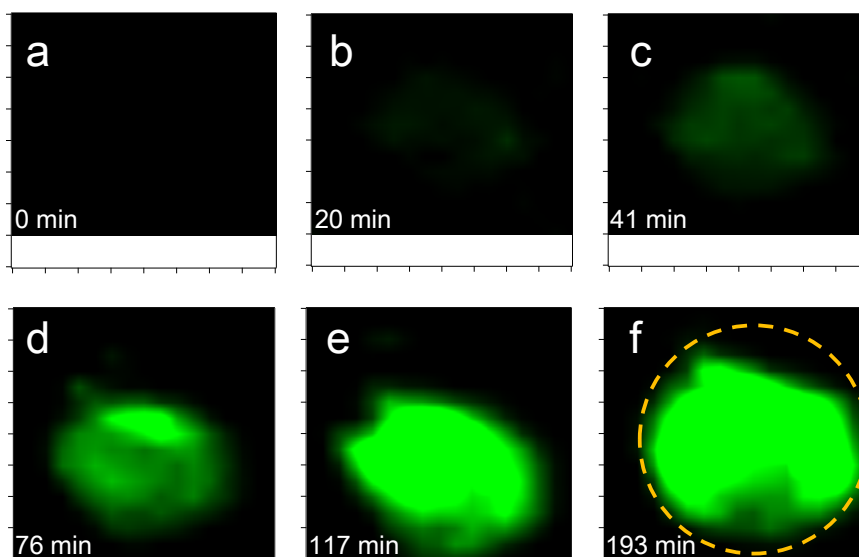


(iii)

497
498

499
500

501
502



503

504

505 (iv) Figure 2. Microscopic image (i) and Raman mapping image of carbonate (ii), nitrate (iii), and sulfate (iv)

506 on the CaCO_3 particle during the multiphase reaction SO_2 with $\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$ on the particle. A-f

507 corresponds to the reaction time of 0, 20, 41, 76, 117, and 193 min. SO_2 : 75 ppm, NO_2 : 75 ppm, RH:

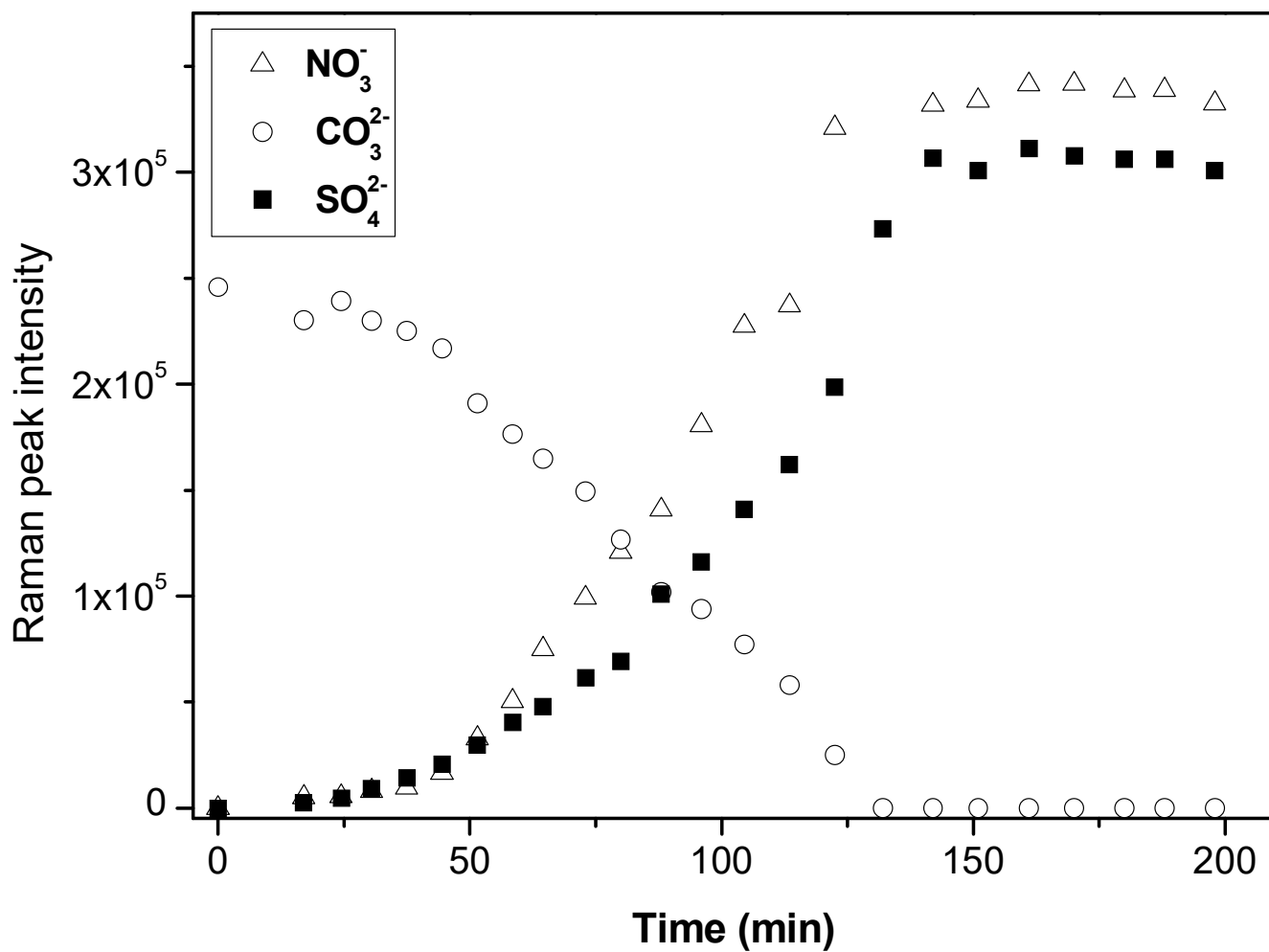
508 72%, O_2 : 20%. The mapping image of carbonate, nitrate, and sulfate are made using the peak area at

509 1050, 1010, and 1087 cm^{-1} , respectively. The red, blue, and green colors indicate the peak intensity of

510 carbonate, nitrate, and sulfate, respectively. The dashed lines in panel iii-f and iv-f indicate the shape of

511 the droplet at the end of the reaction.

512



513
 514 Figure 3. Time series of the Raman peak intensity of NO_3^- , SO_4^{2-} , and CO_3^{2-} during the reaction of SO_2
 515 with $\text{O}_2/\text{NO}_2/\text{H}_2\text{O}$ on CaCO_3 particles. SO_2 : 75 ppm, NO_2 : 75 ppm, RH: 72%, O_2 : 20%. The intensity of
 516 NO_3^- , SO_4^{2-} , and CO_3^{2-} show the peak area at 1050, 1010, and 1087 cm^{-1} , respectively, in Raman spectra
 517 obtained by Raman mapping.

518
 519

1 **Supplement of**
2 **NO₂-initiated multiphase oxidation of SO₂ by O₂ on CaCO₃ particles**

3 Ting Yu^{*}, Defeng Zhao^{*}, Xiaojuan Song, Tong Zhu

4 BIC-ESAT and SKL-ESPC, College of Environmental Sciences and Engineering, Peking University, Beijing, 100871, China

5 ^{*}These authors contributed equally to this work.

6 Correspondence to: Tong Zhu (tzhu@pku.edu.cn)

7

8 S1. Influence of gas phase diffusion on reactive uptake coefficients

9 The Influence of the gas phase diffusion on reactive uptake coefficients was analyzed using the resistor
10 model described by Davidovits et al. (2006) and the references therein.

$$11 \frac{1}{\gamma} = \frac{1}{\Gamma_{diff}} + \frac{1}{\alpha} + \frac{1}{\Gamma_{sat} + \Gamma_{rxn}} \quad (1)$$

12 where Γ_{diff} is the transport coefficient in the gas phase, $1/\Gamma_{diff}$ is the resistance due to the diffusion in the
13 gas phase. Similarly, $1/\Gamma_{sat}$ and $1/\Gamma_{rxn}$ are the resistance due to liquid phase saturation and liquid phase
14 reaction, respectively. α is the mass accommodation coefficient of SO_2 .

15 $1/\Gamma_{diff}$ can be determined using the following equation:

$$16 \frac{1}{\Gamma_{diff}} = \frac{0.75 + 0.238Kn}{Kn(1 + Kn)} \quad (2)$$

17 where Kn is Knudsen number. Knudsen number is defined as

$$18 Kn = \frac{\lambda}{a} \quad (3)$$

19 where λ is the mean free path of a molecule in the gas phase and a is the radius of the particle.

20 λ can be derived from

$$21 \lambda = \frac{3D_g}{c} \quad (4)$$

22 where D_g is the diffusion coefficient in the gas phase and c is the mean molecular velocity.

23 c is determined by

$$24 c = \sqrt{\frac{8RT}{\pi M}} \quad (5)$$

25 where R is the gas constant, T is temperature, and M is the molecular mass of SO_2 .

26 $1/\Gamma_{diff}$ was calculated to be 78 and $1/\gamma$ was calculated to be $\sim 8.3 \times 10^4$. $1/\Gamma_{diff}$ only accounted for $< 0.1\%$
27 of $1/\gamma$. Therefore, the reactive uptake of SO_2 in this study was not limited by gas phase diffusion.

28 The same conclusion can also be drawn by calculating the gas phase diffusion correction factor for a
29 reactive uptake coefficient according to the method in Pöschl et al. (2007) (Equation 20 in their study,
30 also shown as equation 6 below).

$$31 C_g = \frac{1}{1 + \gamma \frac{0.75}{Kn}} \quad (6)$$

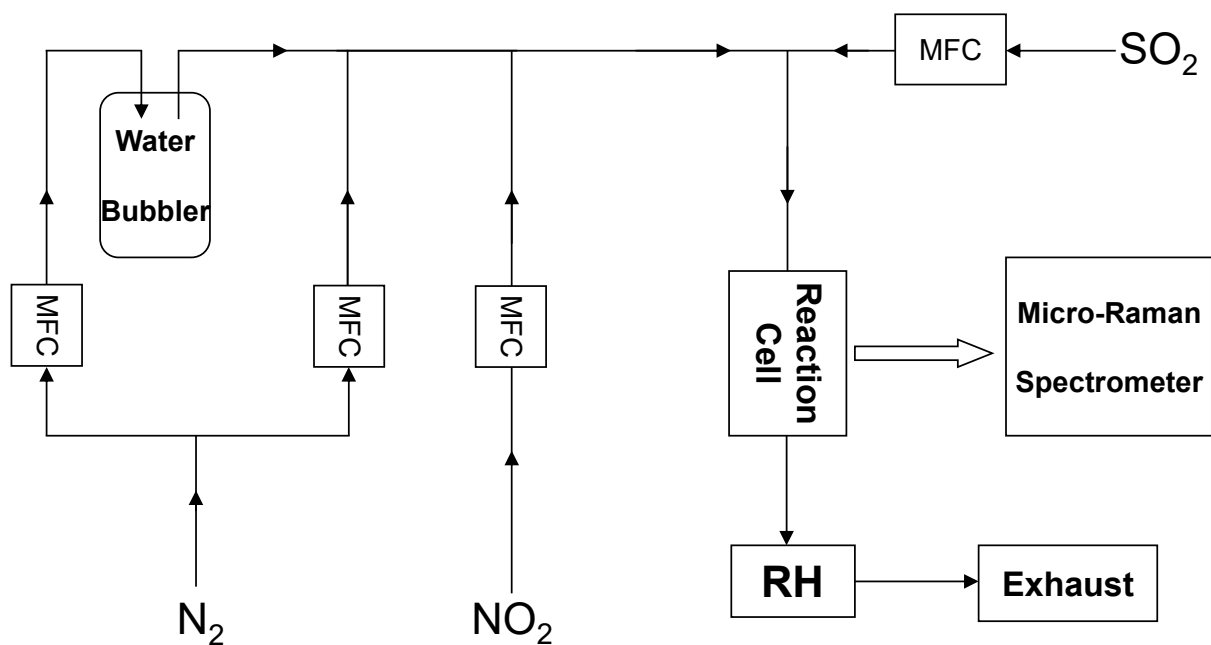
32 where C_g is the gas phase diffusion correction factor for a reactive uptake coefficient.

33 **References**

- 34 Davidovits, P., Kolb, C. E., Williams, L. R., Jayne, J. T., and Worsnop, D. R.: Mass accommodation and
35 chemical reactions at gas-liquid interfaces, *Chem. Rev.*, 106, 1323-1354, 10.1021/cr040366k, 2006.
- 36 Pöschl, U., Rudich, Y., and Ammann, M.: Kinetic model framework for aerosol and cloud surface chemistry and
37 gas-particle interactions - Part 1: General equations, parameters, and terminology, *Atmos. Chem. Phys.*, 7,
38 5989-6023, 2007.
- 39 Zhao, D., Song, X., Zhu, T., Zhang, Z., and Liu, Y.: Multiphase Reaction of SO₂ with NO₂ on CaCO₃ Particles.
40 1. Oxidation of SO₂ by NO₂, *Atmos. Chem. Phys. Discuss.*, 2017, 1-23, 10.5194/acp-2017-610, 2017.

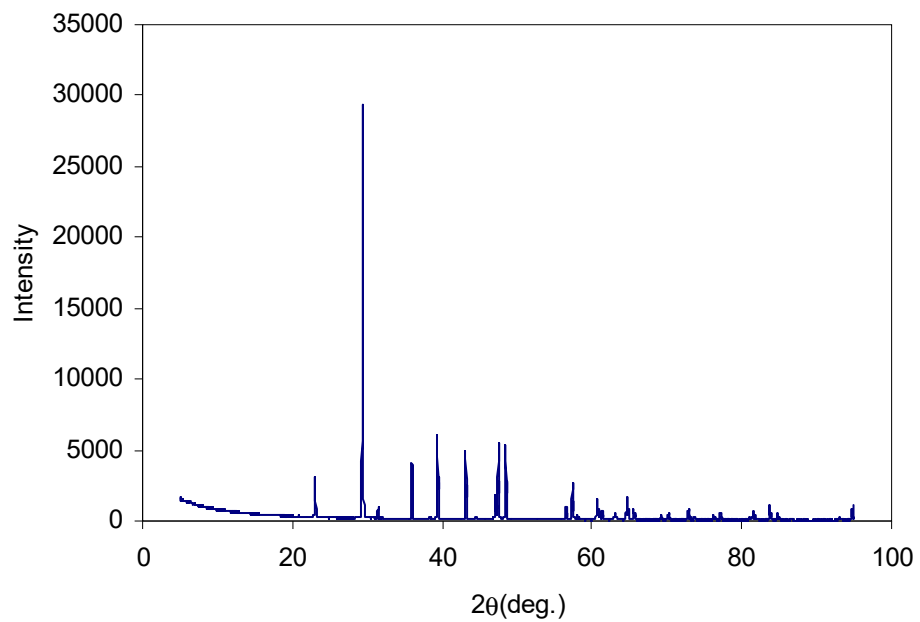
41

42



43

44 Figure S1. Schematic diagram of the experimental setup (Zhao et al., 2017). MFC: mass flow controller.



45

46

47

Figure S2. X-ray diffraction spectra of CaCO₃ particles.