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# Lower than expected volatility of secondary organic aerosols formed during α-pinene ozonolysis

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Abstract. Traditional yield curve analysis shows that semi-volatile organic compounds are a major component of secondary organic aerosols (SOAs). We investigated the volatility distribution of SOAs from α-pinene ozonolysis using positive electrospray ionization mass analysis and dilution- and heat-induced evaporation measurements. Laboratory chamber experiments were conducted on α-pinene ozonolysis, in the presence and absence of OH scavengers. Among these, we identified not only semi-volatile products, but also less volatile highly oxygenated molecules (HOMs) and dimers. Ozonolysis products were further exposed to OH radicals to

- 15 check the effects of photochemical aging. HOMs were also formed during OH-initiated photochemical aging. Most HOMs that formed from ozonolysis and photochemical aging had ten or less carbons. SOA particle evaporation after instantaneous dilution was measured at <1% and ~40% relative humidity. The particle volume fraction remaining decreased with time and stabilised more than 3 h after dilution. The formation of compounds less volatile than of semi-volatile products, and the slow SOA evaporation, contradict the assumptions of current
- 20 atmospheric simulation models.

## 1 Introduction

Atmospheric fine aerosols are believed to negatively affect climate (IPCC, 2013) and human health (Dochery et al., 1993). Secondary organic aerosols (SOAs) are a major component of atmospheric fine aerosols (Zhang et al., 2007). Volatility basis-set (VBS) models have improved the prediction of atmospheric SOA levels by accounting

25 for the decrease in SOA volatility with photochemical aging (Robinson et al., 2007). Our group recently compared the results of VBS model simulations with observational data from several urban and rural sites leeward of mainland East Asia (Morino et al., 2015). The VBS model improved the prediction of ambient organic aerosol levels in spring and summer, but still included large uncertainties.

The volatility distribution of SOAs, a key property in the prediction of particle levels, has been investigated by various techniques. The volatility distribution was evaluated from the chamber results of a SOA yield curve (Lane et al., 2008). This analysis assumes single-generation oxidation and instantaneous gas/particle partitioning for

chamber experiments. Another technique to study particle volatility distribution is heat-induced evaporation. This technique was often applied for SOAs formed from the ozonolysis of  $\alpha$ -pinene, a typical SOA source in the troposphere (Huffman et al, 2009; Epstein et al., 2010; Salo et al., 2011; Kolesar et al., 2015). Heat-induced

35 evaporation provides the volatility at high temperatures, and the enthalpy of vaporization is needed to determine the volatility at ambient temperatures. Furthermore, thermal decomposition may affect results obtained by this method.

Yet another technique to study volatility distribution is dilution-induced evaporation (Grieshop et al., 2007; Vaden et al., 2011; Wilson et al., 2014; Yli-Juuti et al., 2017). Previous studies reported that the time scale for evaporation





40 of  $\alpha$ -pinene SOA after dilution is much longer than that assumed in VBS models. Grieshop et al. (2007) reported that gas/particle partitioning is reversible; however, later research showed that the evaporation process is strongly influenced by the particle phase state, and suggested that dilution results cannot be simply interpreted by gas/particle partitioning.

The chemical analysis of particles can also provide information on the volatility and formation mechanisms of products. Chemical or electrospray ionization mass spectrometry can identify a wide range of oxygenated organic molecules. Using these techniques revealed that the major particle products from α-pinene ozonolysis are highly

- oxygenated molecules (HOMs), ester dimers, and semi-volatile compounds (Zhang et al., 2015; 2017). Recent studies on the parameterization of saturation concentration (Shiraiwa et al., 2014; Li et al., 2016) and the sensitivity of electrospray ionization mass spectrometry (Kruve et al., 2013; Heinritzi et al., 2016) is helpful to
  evaluate SOA volatility using mass analysis data.
- HOMs are believed to be multifunctional peroxides (Ehn et al., 2014) or multifunctional acids (Szmigielski et al., 2007). HOMs would be formed through the auto-oxidation of organic peroxy (RO<sub>2</sub>) radicals (Zhang et al., 2017) or organic oxy (RO) radicals (Müller et al., 2012). Ester dimers identified in α-pinene SOAs might be produced by the dehydration of particles (Yasmeen et al., 2010), reactions of stabilised Criegee intermediates (Kristensen et al.,
- 55 2014; 2016), or gas reactions between two acylperoxy radicals followed by heterogeneous processes (Zhang et al., 2015).

In this study, we performed chamber experiments on  $\alpha$ -pinene ozonolysis in the presence and absence of OH scavengers. Ozonolysis products were further exposed to OH radicals to check the effect of photochemical aging. The volatility distribution of SOAs were evaluated with a wide range of techniques, including offline chemical

60 analysis and dilution- and heat-induced evaporation. We employed positive electrospray ionization mass analysis to detect HOMs and dimers (Zhang et al., 2017). This work aims to evaluate the volatility distribution of α-pinene SOAs using three different techniques and discuss SOA formation processes for improvement of current atmospheric models.

# 2 Methods

## 65 2.1. Chamber experiment

Figure S1 shows a schematic diagram of the laboratory chamber system and analytical instruments used in this study. Initial concentrations are listed with the mass concentration and mean size of the SOA particles produced (Table 1). A 6 m<sup>3</sup> Teflon-coated stainless steel chamber (Sato et al., 2007) was used for experiments under dry conditions (runs 1–9). Prior to each experiment, the chamber was filled with purified air from a custom-made air

- 70 purifier (Horiba Stec Ltd., Japan; relative humidity < 1%). The temperature of the chamber was controlled at 298  $\pm$  1 K. Required amounts of  $\alpha$ -pinene (0.15–0.84 ppmv) and ozone (0.53–1.13 ppmv) were injected into the chamber. In runs 1–4, diethyl ether was also added as an OH scavenger, to suppress secondary reactions with OH radicals from the ozonolysis. In run 6, CO was used as an alternative OH scavenger. In run 9, 1 ppmv methyl nitrite was added as an OH source after  $\alpha$ -pinene was consumed, and the mixture was then irradiated by light from 19 xenon
- 75 lamps (1 kW each); the light was passed through Pyrex filters. The NO<sub>2</sub> photolysis rate was 0.29 min<sup>-1</sup>. The concentrations of α-pinene, ozone, pinonaldehyde (PA), and methyl nitrite were measured every 6 min with an *in-situ* Fourier transform-infrared (FT-IR) spectrometer (Nexus 670, Thermo Fisher Scientific, United States of America (USA)) with a 221.5 m optical path. Experiments under humid conditions (runs 10–12) were performed with a 6 m<sup>3</sup> fluorinated ethylene polyethylene (FEP) film bag (1.81 × 1.81 × 1.81 m, 50 µm). The clean air was
- supplied by purified air generator (DAR-2200, Horiba Ltd., Japan). The relative humidity (RH) measured in the





bag was ~40%. The temperature of the laboratory was controlled at  $298 \pm 1$  K. The ozone concentration in the bag was monitored every 1 min by an ozone monitor (Model 1150, Dylec Inc., Japan). For dilution-induced evaporation measurements, ~20 ppmv CO was added to the reaction mixture as a dilution marker. The CO levels before and after dilution were measured using a CO monitor (Model 48i-TLE, Thermo Fisher Scientific, USA).

85 The coefficient of determination (R<sup>2</sup>) was > 0.99 for the linear least squares analysis of the CO calibration data (3.5 –168 ppmv, n = 8).
 Particle size distribution was observed every 3 min using a scanning mobility particle sizer (SMPS) (Model 3034,

TSI Inc., USA). The effective density of the particles was measured using a custom-made differential mobility analyser (DMA) (Sibata Scientific Technology Ltd., Japan), an aerosol particle mass analyser (APM) (Model 3600,

- 90 Kanomax Inc., Japan), and a condensation particle counter (CPC) (Model 3772, TSI Inc., USA). The effective density of α-pinene ozonolysis SOA was determined to be 1.34 ± 0.12 g cm<sup>-3</sup>.
  A high-resolution time-of-flight aerosol mass spectrometer (AMS) (H-ToF-AMS, Aerodyne Research, USA) (Aiken et al., 2008) combined with a thermodenuder (TD) (ARI thermal denuder, Aerodyne Research, USA) (Faulhaber et al., 2008) was used to measure heat-induced evaporation. The residence time in the heating zone of
- 95 thermodenuder was ~17 s. SOA mass concentration was measured every 3 min using the TD-AMS. Pinonic acid particles generated from an aqueous solution of pinonic acid (13 mM) using an aerosol generator (Model 3076, TSI Inc., USA), and then dried with a diffusion dryer (Model 3062, TSI Inc., USA), were used as a reference for the TD-AMS measurements.

Gas and particle products were measured using a proton transfer reaction-mass spectrometer (PTR-MS)

- 100 (PTR-QMS 500, Ionicon Analytik, Austria) (Lindinger et al., 1998) to determine the saturation concentration of each product. The detailed procedure has been explained elsewhere (Inomata et al., 2014). To measure the products in the gas phase, online measurements were taken from the filtered chamber air after the reaction finished. After this, products in the particle phase were measured; particles were collected on a Fluoropore Teflon filter (Sumitomo Electric Industries, Japan; 47 mmφ, pore size 1 µm) at 16.7 L min<sup>-1</sup> for 30 min. The sample filter was
- 105 placed in a filter holder which was then heated at 358 K under a stream of nitrogen; the gases that evaporated from the filter were measured using the PTR-MS. The saturation concentration was calculated from gas/particle ratio measured by the PTR-MS for each m/z, assuming gas/particle equilibrium.

## 2.2. LC/MS analysis

Offline chemical analysis was performed using the procedure described in Sato et al. (2016). We used positive electrospray ionization-liquid chromatograph/time-of-flight-mass spectrometry (LC/MS). After the reaction finished, particle products were collected on another Teflon filter (Sumitomo Electric Industries, Japan, 47 mmφ, pore size 1 µm) at 16.7 L min<sup>-1</sup> for 30 min. After sampling, the filter sample was sonicated in 5 mL methanol for 30 min using an ultrasonic bath (Model UT-105S, Sharp, Japan; 130 W). The filter extract was concentrated to near dryness under a gentle stream of nitrogen (typically ~1 L min<sup>-1</sup>). A 1 mL formic acid–methanol–water solution

- 115 (v/v/v = 0.05/100/99.95) was added to the concentrated extract to obtain the analytical sample. A 10 μL aliquot of the analytical sample was injected into the LC/MS instrument (LC-TOF, Agilent Technologies, United Kingdom (UK)). The sample was separated with an octadecyl silica gel column (Inertsil ODS-3, GL Science, Japan; 0.5 μm × 3.0 mm × 150 mm). A formic acid–water solution (0.05% v/v) and methanol were used as mobile phases. The total flow of the mobile phases was 0.4 mL min<sup>-1</sup>. The methanol fraction during each analysis was set at 10% (0
- 120 min), 90% (30 min), 90% (40 min), 10% (45 min), and 10% (60 min). The compounds from the column were analysed with the time-of-flight mass spectrometer.





## 2.3. Dilution-induced evaporation experiment

Another 6 m<sup>3</sup> FEP film bag (2.55 × 1.53 × 1.53 m, 50 µm) was used as an external dilution chamber (EDC). Clean air used for the dilution of SOAs was prepared using a clean air generator (Model 1160, Thermo Fisher Scientific, USA) for experiments at RH < 1%, and a purified air generator (DAR-2000, Horiba Stec Ltd., Japan) for experiments at RH ≈ 40%. The temperature of the laboratory was controlled at 298 ± 1 K. Prior to each</li>

- dilution-evaporation experiment, the EDC was filled with clean air. A necessary amount of reaction chamber air was injected into the EDC using a dilution ejector (FPS-4000, Dekati Ltd., Finland). The dilution ratio (DR) was 20–77. Dry filtered air was used as the carrier of the dilution ejector. The particle size distribution, particle density,
- 130 and CO concentration in the EDC were monitored after the SOA dilution.

## 3 Results and Discussion

## 3.1. Time series

Figure 1 shows the time series of gas and particle concentrations and particle oxygen-to-carbon (O/C) ratios measured during the ozonolysis (run 8) and aging (run 9) experiments; these are shown with broken and solid lines,

- 135 respectively. The results measured in run 8 are typical in simple ozonolysis experiments. The concentration of  $\alpha$ -pinene decreased whereas the PA and SOA concentrations increased after the addition of ozone. PA is a gaseous first-generation product from  $\alpha$ -pinene. After most of the  $\alpha$ -pinene was consumed, PA and SOA levels became constant or decreased slowly due to loss from wall deposition in run 8. In run 9, methyl nitrite was added to the reaction mixture 58 min after the start of ozonolysis and it was then irradiated with the xenon lamps. After
- 140 irradiation started, gaseous PA levels decreased and SOA levels increased due to photochemical aging. Methyl nitrite was consumed within 60 min after injection. The total OH exposure for 60 min was estimated to be  $(3.1-3.8) \times 10^{11}$  radical cm<sup>-3</sup> s, based on a similar analysis described in Sato et al. (2016); this corresponds to an OH exposure of 3.6–4.4 d when the OH level is assumed to be  $1.0 \times 10^{6}$  radical cm<sup>-3</sup>.
- The O/C ratio of SOA was ~0.40 immediately after SOA concentrations increased above detectable levels, and then decreased with time. The O/C ratio decreased to ~0.34 when most of the α-pinene was consumed and then increased slowly in run 8. Saturated organic compounds are formed from α-pinene ozonolysis; therefore, ozone-initiated reactions of the secondary products must be slow. In run 9, the O/C ratio increased to ~0.44 as a result of OH exposure, suggesting that highly oxygenated SOAs formed from photochemical aging. The results suggest that the reactions between the gaseous secondary products and OH radicals will be a major source of SOA
- 150 formation during photochemical aging.

## 3.2. Chemical composition analysis

Strong signals measured during positive electrospray ionization analysis were identified as sodium-attached product ions. Although several signals were identified as protonated product ions, they were weaker than those of sodium-attached ions. Table S1 shows monomer ( $C_{8-10}H_{12-18}O_{3-9}$ ) and dimer ( $C_{16-20}H_{24-32}O_{4-12}$ ) products identified

as sodium-attached ions. Figure 2 shows the extracted ion chromatograms (mass extraction window  $\pm 20$  ppm) measured for  $C_{10}H_{16}O_5Na^+$ ,  $C_{10}H_{16}O_6Na^+$ , and  $C_{10}H_{16}O_7Na^+$ . No sodium salt was added to the mobile phase or analytical sample. The signal intensities of the sodium-attached ions were confirmed to have a linear relationship with the amount of injected sample.

 $C_{10}H_{16}O_{5-7}$  compounds are HOMs with the same number of carbon and hydrogen atoms as pinonic acid. We also detected sodium-attached ions of pinonic acid ( $C_{10}H_{16}O_3$ ) and its hydroxylated derivative ( $C_{10}H_{16}O_4$ ), both of





which are typical  $\alpha$ -pinene SOA markers. The carbon oxidation states ( $\approx 2 \text{ O/C} - \text{H/C}$ ; Kroll et al., 2011) of  $C_{10}H_{16}O_3$ ,  $C_{10}H_{16}O_4$ ,  $C_{10}H_{16}O_5$ ,  $C_{10}H_{16}O_6$ , and  $C_{10}H_{16}O_7$  are -1.0, -0.8, -0.6, -0.4, and -0.2, respectively. In this study, we define HOMs as molecules with carbon oxidation states of -0.6 or higher. Several peaks were observed in the chromatogram for each HOM, suggesting that several isomers are available. New peaks appeared, or peak

165 heights increased, in chromatograms from the photochemical aging experiment (run 9; red lines) compared to the simple ozonolysis experiment (run 8; blue lines), suggesting that HOMs are produced during photochemical aging as well as ozonolysis.

Figure 3 shows the carbon number distributions measured in selected experiments; runs 1 and 6 are experiments with OH scavengers; runs 7 and 8 are experiments without OH scavengers; and run 9 is the photochemical aging

- 170 experiment. We employed column-separated results to determine the distribution of products; this enabled us to remove any contributions from cluster signals originating from the mobile phases. The total area of the chromatographic peaks were calculated over retention times of 4–40 min to determine the signal abundance. Baseline signals and peaks appearing at short (< 3 min) retention times were excluded. Signal abundances were calculated for the products listed in Table S1. The abundance calculated for each product was then normalized</p>
- 175 using the total abundance of the measured products. The calculated results reflected the relative abundances of particle products, if decomposition of products during sampling and pre-treatment was negligible. Nine- and ten-carbon products with carbon oxidation states less than -0.6 were semi-volatile products such as pinic acid and pinonic acid. These products were detected in all selected runs. Eight- to ten-carbon HOMs were also detected in all runs. The highest relative abundance was observed in the photochemical aging experiment,
- followed in order by the experiment without OH scavengers and the experiment with OH scavengers. These results suggest that HOMs are formed by the OH-initiated aging of  $\alpha$ -pinene oxidation products or OH-initiated oxidation of  $\alpha$ -pinene. HOMs were detected in the presence of OH scavengers, suggesting that HOMs are also formed from  $\alpha$ -pinene ozonolysis. The auto-oxidation of RO<sub>2</sub> radicals (Zhang et al., 2017) or RO radicals (Müller et al., 2012) could explain present results. Both types of radicals are formed from the ozonolysis of  $\alpha$ -pinene as well as the
- 185 OH-initiated oxidation of α-pinene and its secondary products. The RO<sub>2</sub> and RO radicals produced undergo intramolecular hydrogen atom shifts, followed by successive oxidation, to form HOMs. Dimers were detected in runs 1 and 6, suggesting that they are formed from the ozonolysis of α-pinene. Most dimers had carbon oxidation states less than -0.6; in other words, dimers were less oxygenated than monomers. If dimers are formed from the simple addition of or dehydration from monomers, their carbon oxidation states must
- 190 be close to those of the precursor monomers. The results suggest that dimers will be formed from less-oxygenated monomers instead of HOMs, and that such dimerization will occur prior to the formation of the latter. The results also suggest that dimers will barely be oxidized during photochemical aging, probably due to their low to extremely low volatility. Stabilised Criegee intermediates and acyl peroxy radicals are not highly functionalised because they are formed directly from α-pinene ozonolysis. Dimers may be formed from the reactions of stabilised
- 195 Criegee intermediates with compounds less oxygenated than HOMs (Kristensen et al., 2014; 2016), or from the recombination of acyl peroxy radicals (Zhang et al., 2015). C<sub>19</sub> and C<sub>20</sub> dimers had the lowest abundances in run 9. These dimers may decompose during photochemical aging although the mechanism is unclear. Previous studies have reported the presence of terpene-derived HOMs and dimers in ambient forest particles (Ehn et al., 2014; Yasmeen et al., 2010); however, dimers are not always detected in association with pinonic acid in
- 200 ambient particles (Koutchev et al., 2014). Therefore, the formation mechanisms of HOMs and dimers from  $\alpha$ -pinene oxidation require further clarification.





#### 3.3. Volatility distribution from chemical analysis

To evaluate SOA volatility distribution from LC/MS results, the saturation concentration ( $C^*$ ) was calculated using the two previous parameterization techniques. Shiraiwa et al. (2014) parameterized molecular weight (MW) as a function of log<sub>10</sub>  $C^*$  (1D function). The 1D function was fitted to saturation concentration data calculated for  $\alpha$ -pinene oxidation products using the EVAPORATION (Estimation of VApour Pressure of ORganics, Accounting for Temperature, Intramolecular, and Non-additivity effects) model (Compernolle et al., 2011). However, Li et al. (2016) parameterized log<sub>10</sub>  $C^*$  as a function of the number of carbon and oxygen atoms (2D

function). The 2D function was fitted to EVAPORATION data for various oxygenated organic compounds. Figure

210 S2 plots the MW of α-pinene oxidation products as a function of  $\log_{10} C^*$ , and compares results obtained from the present PTR-MS measurements with the results calculated using EVAPORATION (Shiraiwa et al., 2014). The fit result of the 1D function is shown as a straight line. Figure 2S also shows the results calculated by Sato et al. (2016) using the SPARC online calculator (Hilal et al., 2003). Similar results were obtained from the EVPORATION model and the SPARC online calculator, both of which are based on modified structure activity

215 relationships. The present PTR-MS results support the predictions of the EVAPORATION model and SPARC online calculator.

The sensitivity of the electrospray ionization mass spectrometer is determined by ionization and transmission efficiencies. The ionization efficiency for sodium adduct formation has a weak correlation with the partition coefficient between vacuum and solvent, but its parameterization will not be simple due to the competition with

- 220 protonated ion formation (Kruve et al., 2013). We assumed that the sodium adduct formation efficiency is constant against MW because both low- and high-MW compounds present in SOAs are mixtures of various oxygenated hydrocarbons. We assumed that the transmission efficiency of the time-of-flight mass spectrometer is proportional to  $(m/z)^{1/2}$  in the region of m/z 200–500.
- Figure 4a shows the volatility distributions of gas plus particle products determined from LC/MS data using a 1D 225 function. The signal intensity of the particle products was corrected, accounting for transmission efficiency as described above. For comparison with conventional SOA volatility data, the abundance of gas products was added to that of particle products. The former was determined based on a gas/particle partitioning model. The gas component of  $\log_{10} C^* = 3$  had an abundance comparable with the particle components, whereas that of  $\log_{10} C^* < 3$  had negligible abundances. The gas plus particle abundances were then normalized using total abundance. All
- 230 plotted results had similar bimodal distributions consisting of semi-volatile organic compounds (with  $\log_{10} C^* = 0$ -3) and lower volatility organic compounds (with  $\log_{10} C^* < 0$ ). The volatility distribution of the semi-volatile component is similar to that suggested from yield curve analysis (Lane et al., 2008), however the lower volatility component has not been predicted by traditional analysis. Organic compounds with  $\log_{10} C^* < 0$  were formed even in the presence of OH scavenger, suggesting that these compounds were formed as the first-generation
- 235 products from  $\alpha$ -pinene ozonolysis. Overall volatility distribution barely changed during OH-initiated aging, although HOMs increased; this suggests that dimers determined the overall volatility distribution of  $\alpha$ -pinene SOAs.

Figure 4b shows the results obtained using a 2D function. The volatility distributions determined by the 2D function shifted to a higher saturation concentration than those determined by the 1D function. Table S2 shows the

240  $\log_{10} C^*$  determined for  $\alpha$ -pinene oxidation products using the SPARC online calculator and 1D and 2D functions. The results predicted for dimers by the 1D function agreed with the SPARC results within an error of  $\pm 2$ ; in contrast, the 2D function predicted values five to six units higher than the SPARC results. The accuracy of the results predicted by the 2D function was probably worse than those predicted by 1D function because the former





was adjusted to oxygenated organic compounds with higher volatilities than α-pinene oxidation products. Better 245 parameterization methods for oxygenated organic compounds are necessary to obtain more accurate volatility distributions.

# 3.4. Heat- and dilution-induced evaporation

Figure S3 shows thermograms measured during heat-induced evaporation measurements of SOAs and pinonic acid particles, where the thermogram is the plot of particle mass fraction remaining (MFR) as a function of

- 250 thermodenuder temperature. The thermograms measured for α-pinene SOAs in this study were very close to previous results (e.g., Salo et al., 2011; Kolesar et al., 2015). The data obtained at each temperature were distributed among five runs with OH scavengers (runs 1, 2, 3, 4, and 6; mean particle diameter = 236–291 nm). Although the effects of OH scavengers, photochemical aging, and relative humidity on particle volatility were studied, the thermograms showed that all SOA results were similar to each other, within experimental uncertainties.
- 255 The compounds comprising SOA particles had lower volatilities than pinonic acid where the log<sub>10</sub> C<sup>\*</sup> of pinonic acid was determined as 2.25 by the SPARC online calculator. Figure 4c shows the volatility distributions determined for gas plus particle products using thermogram data. Thermogram data were converted into volatility distributions using a previous empirical method developed by Faulhaber et al. (2009). In that study, the relationship between saturation concentration and the temperature at
- 260 which MFR = 0.5 was measured directly for several kinds of single compound particles. The present study, and previous ones, used similar residence times in the heating zone of a thermodenuder. The volatility distributions determined under various conditions were again similar to each other within experimental uncertainties. The results of heat-induced evaporation also suggest that there is a substantial low or extremely low volatility component in addition to the semi-volatile component.
- Figure 5 shows the time series of the volume fraction remaining (VFR) measured for SOA particles after dilution. The number concentration and mean size of the SOA particles decreased with time. We assumed that decrease in particle numbers resulted from loss by deposition on the wall, and that the decrease in the mean diameter of the particles resulted from evaporation. To remove the influence of wall loss, VFR was determined by  $r^3/r_0^3$  where *r* is the geometrical mean size of particles and  $r_0$  is the geometrical mean size immediately before dilution. The red
- 270 symbols show results obtained at a dilution ratio (DR) of 20–75 under dry conditions. Although evaporation is assumed to occur instantaneously in VBS models, the VFR of SOA particles decreased very slowly and became stable more than 3 h after dilution. These results support the results reported by previous researchers (e.g. Grieshop et al., 2007; Vaden et al., 2011). Straight red lines represent equilibrium levels calculated based on previous yield curve data. The so-called 'base case' yield curve reported by Lane et al. (2008) was used for the calculations. The
- 275 VFR of SOA particles measured 3 h after dilution was higher than equilibrium levels based on yield curve data. Mean particle size may also have decreased due to size-dependent wall loss. The decrease in the VFR resulting from size-dependent wall loss was calculated using directly measured size-dependent wall loss rate, and was determined to be ~2% for 3 h of measurement. The evaluated decrease in the VFR was less than the statistical error of the SMPS volume (~5%). The blue symbols are the results obtained under humid conditions. The rate of
- 280 evaporation depended strongly on RH due to the changes in the viscosity or chemical composition of particles (Wilson et al., 2014; Yli-Juuti et al., 2017). The VFR measured under humid conditions also decreased to levels higher than equilibrium levels based on yield curve (blue lines). The present results indicate that gas/particle partitioning was virtually irreversible even though the VFR continued to decrease after 3 h. The present results based on SOA particles contrast with previous results based on semi-volatile diesel particles, which evaporate





285 instantaneously after dilution (Fujitani et al., 2012). These results strongly suggest that SOA particles include low or extremely low volatility organic compounds.

#### 4 Implication and Conclusions

In this study, we investigated the effect of OH scavenger and photochemical aging on the molecular distribution of α-pinene ozonolysis products including HOMs and dimers. By employing positive electrospray ionization combined with recent parameterization techniques for saturation concentration, we also studied product volatility distribution. To the best of our knowledge, this study is the first to arrive at the results described above.

The first-generation products formed during  $\alpha$ -pinene ozonolysis were found to include compounds less volatile than semi-volatile compounds, and dilution-induced evaporation was found to be much slower than that assumed in the VBS model. These findings are important, and they contradict the current VBS model, in which

295 direct formation of semi-volatile organic compounds and instantaneous dilution-induced evaporation are assumed. If dilution-induced evaporation is slower than that expected by the current VBS model, ambient SOA level after dilution would be underestimated by the model, and photochemical aging in gas phase would be slower than that expected in the model. In this study, no clear decrease in SOA volatility during photochemical aging could be observed; this may be explained by the slow evaporation of SOA particles. Further improvement of the 300 atmospheric simulation model will be necessary to explain both laboratory and ambient SOA levels.

The products of  $\alpha$ -pinene ozonolysis were semivolatile compounds, HOMs, and dimers, and HOMs also formed during OH-initiated photochemical aging. Dimers were less oxygenated than monomers. These results are consistent with previously proposed mechanisms in which the formation of dimers is initiated by reactions of stabilized Criegee intermediates or acylperoxy radicals, and the formation of HOMs is initiated by auto-oxidation

305 of RO or RO<sub>2</sub> radicals.

*Author contribution.* K. S., S. R., and T. I. designed and performed chamber experiments and also conducted TD-AMS and LC/MS analyses. Y. F., A. F., and Y. K. designed and carried out dilution measurements. S. I. and H. T. designed and carried out PTR-MS measurements. Y. M., K. T., and S. S. contributed to data interpretations from

310 the viewpoint of atmospheric modelling. T. H., A. S., and A. T. gave technical support for TD-AMS measurements and also contributed to data interpretations.

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## References

320 Aiken, A. C., DeCarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S., Ulbrich, I. M., Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q., Trimborn, A., Northway, M., Ziemann, P. J., Canagaratna, M. R., Onasch, T. B., Alfarra, M. R., Prevot, A. S. H., Dommen, J., Duplissy, J., Metzger, A., Baltensperger, U., and





330

Jimenez, J. L.: O/C and OM/OC ratios of primary, secondary, and ambient organic aerosols with high-resolution time-of-flight aerosol mass spectrometry, Environ. Sci. Technol., 42, 4478–4485, doi:10.1021/es703009q, 2008.

325 Compernolle, S., Ceulemans, K., and Müller, J. F.: EVAPORATION: a new vapour pressure estimation method for organic molecules including non-additivity and intramolecular interactions, Atmos. Chem. Phys., 11, 9431– 9450, doi:10.5194/acp-11-9431-2011, 2011.

Dockery, D. W., Pope, A., III, Xu, X., Spengler, J. D., Ware, J. H., Fay, M. E., Ferris, B. G., Jr., and Speizer, F. E.: An association between air pollution and mortality in six U.S. cities, N. Engl. J. Med., 329, 1753–1759, doi:10.1056/NEJM199312093292401, 1993.

Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R.,
Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma, J.,
Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., Canagaratna, M., Dal
Maso, M., Berndt, T., Petäjä, T., Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J., and

- Mentel, T. F.: A large source of low-volatility secondary organic aerosol, Nature, 506, 476–479, doi:10.1038/nature13032, 2014.
  Epstein, S. A., Riipinen, I., and Donahue, N. M.: A semiempirical correlation between enthalpy of vaporization and saturation concentration for organic aerosol, Environ. Sci. Technol., 44, 743–748, doi:10.1021/es902497z, 2010.
- 340 Faulhaber, A. E., Thomas, B. M., Jimenez, J. L., Jayne, J. T., Worsnop, D. R., and Ziemann, P. J.: Characterization of a thermodenuder-particle beam mass spectrometer system for the study of organic aerosol volatility and composition, Atmos. Meas. Tech., 2, 15–31, doi:10.5194/amt-2-15-2009, 2009. Fujitani, Y., Saitoh, K., Fushimi, A., Takahashi, K., Hasegawa, S., Tanabe, K., Kobayashi, S., Furuyama, A., Hirano, S., and Takami, A.: Effect of isothermal dilution on emission factors of organic carbon and n-alkanes in
- 345 the particle and gas phases of diesel exhaust, Atmos. Environ., 59, 389–397, doi:10.1016/j.atmosenv.2012.06.010, 2012.

Grieshop, A. P., Donahue, N. M., and Robinson, A. L.: Is the gas-particle partitioning in alpha-pinene secondary organic aerosol reversible?, Geophys. Res. Lett., 34, L14810, doi:10.1029/2007GL029987, 2007.

Heinritzi, M., Simon, M., Steiner, G., Wagner, A. C., Kürten, A., Hansel, A., and Curtius, J.: Characterization of
the mass-dependent transmission efficiency of a CIMS, Atmos. Meas. Tech., 9, 1449–1460,
doi:10.5194/amt-9-1449-2016, 2016.

Hilal, S. H., Karickhoff, S. W., and Carreira, L. A.: Prediction of the vapor pressure, boiling point, heat of vaporization and diffusion coefficient of organic compounds, QSAR Comb. Sci., 22, 565–574, doi:10.1002/qsar.200330812, 2003.

- 355 Huffman, J. A., Docherty, K. S., Mohr, C., Cubison, M. J., Ulbrich, I. M., Ziemann, P. J., Onash, T. B., and Jimenez, J. L.: Chemically-resolved volatility measurements of organic aerosol from different sources, Environ. Sci. Technol., 43, 5351–5357, doi:10.1021/es803539d, 2009. Inomata, S., Sato, K., Hirokawa, J., Sakamoto, Y., Tanimoto, H., Okumura, M., Tohno, S., and Imamura, T.: Analysis of secondary organic aerosols from ozonolysis of isoprene by proton transfer reaction mass spectrometry,
- 360 Atmos. Environ., 97, 397–405, doi:10.1016/j.atmosenv.2014.03.045, 2014.
  IPCC: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, edited by: Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M., Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 1535 pp., 2013.





365 Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R., Wilson, K. R., Altieri, K. E., Mazzoleni, L. R., Wozniak, A. S., Bluhm, H., Mysak, E. R., Smith, J. D., Kolb, C. E., and Worsnop, D. R.: Carbon oxidation state as a metric for describing the chemistry of atmospheric organic aerosol, Nat. Chem., 3, 133–139, doi:10.1038/NCHEM.948, 2011.

Kolesar, K. R., Li, Z., Wilson, K. R., and Cappa, C. D.: Heating-induced evaporation of nine different secondary
organic aerosol types, Environ. Sci. Technol., 49, 12242–12252, doi:10.1021/acs.est.5b03038, 2015.

Koutchev, I., Fuller, S. J., Giorio, C., Healy, R. M., Wilson, E., O'Connor, I., Wenger, J. C., McLeod, M., Aalto, J., Ruuskanen, T. M., Maenhaut, W., Jones, R., Venables, D. S., Sodeau, J. R., Kulmala, M., and Kalberer, M.: Molecular composition of biogenic secondary organic aerosols using ultrahigh-resolution mass spectrometry: comparing laboratory and field studies, Atmos. Chem. Phys., 14, 2155–2167, doi:10.5194/acp-14-2155-2014, 2014.

- Kristensen, K., Cui, T., Zhang, H., Gold, A., Glasius, M., and Suratt, J. D.: Dimers in  $\alpha$ -pinene secondary organic aerosol: effect of hydroxyl radical, ozone, relative humidity and aerosol acidity, Atmos. Chem. Phys., 14, 4201–4218, doi:10.5194/acp-14-4201-2014, 2014.
- Kristensen, K., Watne, Å. K., Hammes, J., Lutz, A., Petäjä, T., Hallquist, M., Bilde, M., Glasius, M.:
  High-molecular weight dimer esters are major products in aerosols from α-pinene ozonolysis and the boreal forest, Environ. Sci. Technol., Lett. 3, 280–285, 10.1021/acs.estlett.6b00152, 2016.
  Kruve, A., Kaupmees, K., Liigand, J., Oss, M., and Leito, I.: Sodium adduct formation efficiency in ESI source, J. Mass Spectrom., 48, 695–702, doi:10.1002/jms.3218, 2013.
- Lane, T. E., Donahue, N. M., and Pandis, S. N.: Simulating secondary organic aerosol formation using the volatility basis-set approach in a chemical transport model, Atmos. Environ., 42, 7439–7451, doi:10.1016/j.atmosenv.2008.06.026, 2008.
  Li, Y., Pöschl, U., and Shiraiwa, M: Molecular corridors and parameterizations of volatility in the chemical evolution of organic aerosols, Atmos. Chem. Phys., 16, 3327–3344, doi:10.5194/acp-16-3327-2016, 2016.
  Lindinger, W., Hansel, A., and Jordan, A.: Proton-transfer-reaction mass spectrometry (PTR-MS): on-line
- monitoring of volatile organic compounds at pptv levels, Chem. Soc. Rev., 27, 347–354, doi:10.1016/S0168-1176(97)00281-4, 1998.
  Morino, Y., Nagashima, T., Sugata, S., Sato, K., Tanabe, K., Noguchi, T., Takami, A., Tanimoto, H., and Ohara, T.: Verification of chemical transport models for PM<sub>2.5</sub> chemical composition using simultaneous measurement data over Japan, Aeros. Air Qual. Res., 15, 2009–2023, doi:10.4209/aagr.2015.02.0120 2015.
- 395 Müller, L., Reinnig, M.-C., Naumann, K. H., Saathoff, H., Mentel, T. F., Donahue, N. M., and Hoffmann, T.: Formation of 3-methyl-1,2,3-butanetricarboxylic acid via gas phase oxidation of pinonic acid – a mass spectrometric study of SOA aging, Atmos. Chem. Phys., 12, 1483–1496, doi:10.5194/acp-12-1483-2012, 2012. Robinson, A. L., Donahue, N. M., Shrivvastava, M. K., Weitkamp, E. A., Sage, A. M., Grieshop, A. P., Lane, T. E., Pierce, J. R., and Pandis, S. N.: Rethinking organic aerosols: Semivolatile emissions and photochemical aging,
- Science, 315, 1259–1262, doi:10.1126/science.1133061 2007.
  Salo, K., Hallquist, M., Jonsson, Å. M., Saathoff, H., Naumann, K.-H., Spindler, C., Tillmann, R., Fuchs, H., Bohn, B., Rubach, F., Mentel, T. F., Müller, L., Reinnig, M., Hoffmann, T., and Donahue, N. M.: Volatility of secondary organic aerosol during OH radical induced aging, Atmos. Chem. Phys., 11, 11055–11067, doi:10.5194/acp-11-11055-2011, 2011.
- 405 Sato, K., Hatakeyama, S., and Imamura, T.: Secondary organic aerosol formation during the photooxidation of toluene: NO<sub>x</sub> dependence of chemical composition, J. Phys. Chem. A, 111, 9796–9808, doi:10.1021/jp071419f, 2007.





410

Sato, K., Jia, T., Tanabe, K., Morino, Y., Kajii, Y., and Imamura, T.: Terpenylic acid and nine-carbon multifunctional compounds formed during the aging of b-pinene ozonolysis secondary organic aerosol, Atmos. Environ., 130, 127–135, doi:10.1016/j.atmosenv.2015.08.047, 2016.

Shiraiwa, M., Berkemeier, T., Schilling-Fahnestock, K. A., Seinfeld, J. H., and Pöschl, U.: Molecular corridors and kinetic regimes in the multiphase chemical evolution of secondary organic aerosol, Atmos. Chem. Phys., 14, 8323–8341, doi:10.5194/acp-14-8323-2014, 2014.

Szmigielski, R., Surratt, J. D., Gómez-González, Y., Van der Veken, P., Kourtchev, I., Vermeylen, R., Blockhuys,

415 F., Jaoui, M., Kleindienst, T. E., Lewandowski, M., Offenberg, J. H., Edney, E. O., Seinfeld, J. H., Maenhaut, W., and Claeys, M.: 3-methyl-1,2,3-butanetricarboxylic acid: An atmospheric tracer for terpene secondary organic aerosol, Geophys. Res. Lett., 34, L24811, doi:10.1029/2007GL031338, 2007.

Vaden, T. D., Imre, D., Beránek, J., Shrivastava, M., and Zelenyuk, A.: Evaporation kinetics and phase of laboratory and ambient secondary organic aerosol, Proc. Natl. Acad. Sci. USA, 108, 2190–2195, doi:10.1073/pnas.1013391108, 2011.

Wilson, J., Imre, D., Beránek, Shrivastava, and Zelenyuk, A.: Evaporation kinetics of laboratory-generated secondary organic aerosols at elevated relative humidity, Environ. Sci. Technol., 49, 243–249, doi:10.1021/es505331d, 2014.

Yasmeen, F., Vermeylen, R., Szmigielski, R., Iinuma, Y., Böge, O., Herrmann, H., Maenhaut, W., and Claeys, M.:

- Terpenylic acid and related compounds: precursors for dimers in secondary organic aerosol from the ozonolysis of α- and β-pinene, Atmos. Chem. Phys., 10, 9383–9392, doi:10.5194/acp-10-9383-2010, 2010.
  Yli-Juuti, T., Pajunoja, A., Tikkanen, O.-P., Buchholz, A., Faiola, C., Väisänen, O., Hao, L., Kari, E., Peräkylä, O., Garmash, O., Shiraiwa, M., Ehn, M., Lehtinen, K., and Virtanen, A.: Factors controlling the evaporation of secondary organic aerosol from α-pinene ozonolysis, Geophys. Res. Lett., 44, 2562–2570,
- 430 doi:10.1002/2016GL072364, 2017. Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I., Alfarra, M. R., Takami, A., Middlebrook, A. M., Sun, Y. L., Dzepina, K., Dunlea, E., Docherty, K., DeCarlo, P. F., Salcedo, D., Onasch, T., Jayne, J. T., Miyoshi, T., Shimono, A., Hatakeyama, S., Takegawa, N., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Willliams, P., Bower, K., Bahreini, R., Cottrell, L., Griffin, R. J.,
- Rautiainen, J., Sun, J. Y., Zhang, Y. M., and Worsnop, D. R.: Ubiquity and dominance of oxygenated species in organic aerosols in anthropogenically-influenced Northern Hemisphere midlatitudes, Geophys. Res. Lett., 34, L13801, doi:10.1029/2007GL029979, 2007.
  Zhang, X., McVay, R., Huang, D. D., Dalleska, N. F., Aumont, B., and Flagan, R. C., Seinfeld, J. H.: Formation

and evolution of molecular products in α-pinene secondary organic aerosol, Proc. Natl. Acad. Sci. USA, 112,
14168–14173, doi:10.1073/pnas.1517742112 2015, 2015.

Zhang, X., Lambe, A. T., Upshur, M. A., Brooks, W., Gray Be, A., Thomson, R. J., Geiger, F. M., Surratt, J. D.,
Zhang, Z., Gold, A., Graf, S., Cubison, M. J., Groessl, M., Jayne, J. T., Worsnop, D. R., and Canagaratna, M. R.:
Highly oxygenated multifunctional compounds in α-pinene secondary organic aerosol, Environ. Sci. Technol., 51,
5932–5940, doi:10.1021/acs.est.6b06588 2017.





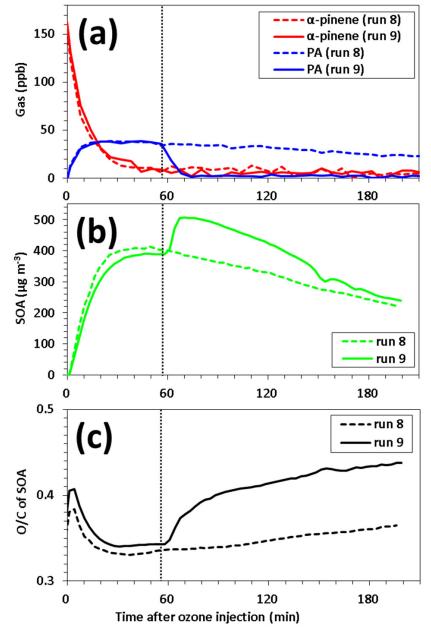
run	reaction <sup>a</sup>	RH %	[HC]₀ ppmv	[O₃]₀ ppmv	[SOA] <sup>c</sup> µg m⁻³	size nm	measurements <sup>d</sup>
2	α-pinene/O <sub>3</sub> /Et <sub>2</sub> O	<1	0.54	1.02	1,539	262	TD-AMS, EDC (DR = 20)
3	$\alpha$ -pinene/O <sub>3</sub> /Et <sub>2</sub> O	<1	0.51	1.08	1,382	236	TD-AMS, EDC (DR = 43)
4	$\alpha$ -pinene/O <sub>3</sub> /Et <sub>2</sub> O	<1	0.53	1.13	1,490	291	TD-AMS
5	$\alpha$ -pinene/O <sub>3</sub>	<1	0.11	0.58	216	198	TD-AMS, PTR-MS
6	$\alpha$ -pinene/O <sub>3</sub> /CO	<1	0.31	0.65	964	257	TD-AMS, LC/MS, PTR-MS
7	$\alpha$ -pinene/O <sub>3</sub>	<1	0.32	0.59	859	274	TD-AMS, LC/MS, PTR-MS
8	$\alpha$ -pinene/O <sub>3</sub>	<1	0.15	0.62	303	226	TD-AMS, LC/MS, PTR-MS
9	$\alpha$ -pinene/O <sub>3</sub> (+ aging <sup>b</sup> )	<1	0.16	0.53	461	233	TD-AMS, LC/MS, PTR-MS
10	α-pinene/O <sub>3</sub> /Et <sub>2</sub> O	~40	0.46	1.09	1,593	259	EDC (DR = 33)
11	α-pinene/O <sub>3</sub> /Et <sub>2</sub> O	~40	0.46	1.09	1,947	298	TD-AMS, EDC (DR = 20)
12	α-pinene/O <sub>3</sub> /Et <sub>2</sub> O	~40	0.46	1.09	1,671	260	EDC (DR = 77)

445 Table 1. Experimental conditions, particle mass concentration, and mean particle size.

<sup>a</sup> Diethyl ether (Et<sub>2</sub>O) and carbon monoxide (CO) were used as OH scavengers. <sup>b</sup> SOA formed from ozonolysis was exposed to OH radicals. <sup>c</sup> Calculated by assuming particle density to 1.34 g cm<sup>-3</sup> (present study). <sup>d</sup> EDC is external dilution chamber; DR is dilution ratio.





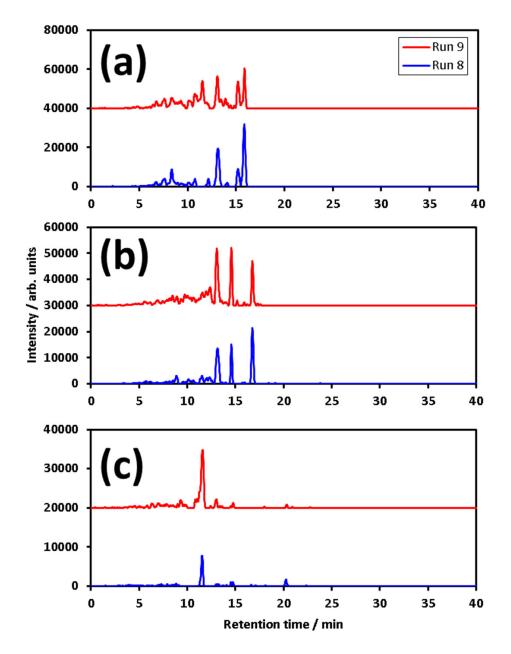


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Figure 1: Time series of (a) FT-IR gas concentrations and (b) AMS SOA concentration and (c) AMS SOA O/C ratio during simple ozonolysis experiment (run 8) and aging experiment, in which SOA formed from ozonolysis was exposed to OH radicals (run 9).







455 Figure 2: Extracted ion chromatograms of (a) C<sub>10</sub>H<sub>16</sub>O<sub>5</sub>Na<sup>+</sup>, (b) C<sub>10</sub>H<sub>16</sub>O<sub>6</sub>Na<sup>+</sup>, and (c) C<sub>10</sub>H<sub>16</sub>O<sub>7</sub>Na<sup>+</sup> formed by the addition of sodium ion during the electrospray ionization of highly oxygenated molecules present in fresh (run 8) and aged SOA particles (run 9).





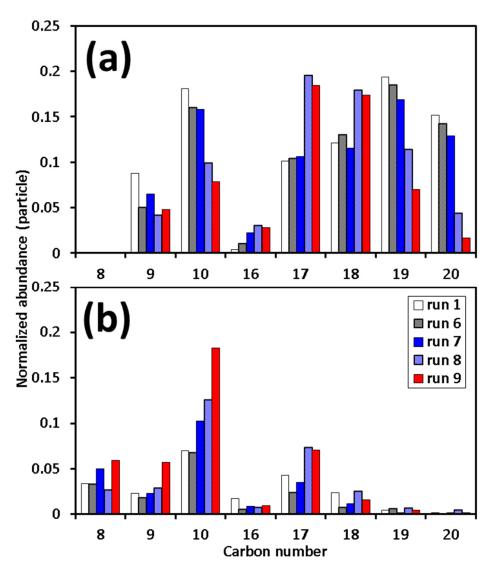


Figure 3: Carbon number distributions determined for particle products at (a) carbon oxidation state < -0.6 (less oxygenated than HOMs) and (b) carbon oxidation state ≥ -0.6 (HOMs); abundance was directly calculated from the summation of signal intensities.





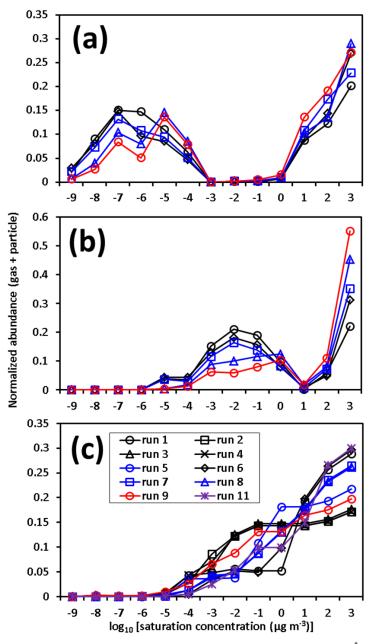


Figure 4: Volatility distributions determined for gas + particle products from LC/MS data using C<sup>\*</sup> parameterization methods by (a) Shiraiwa et al. (2014), (b) Li et al. (2016) and (c) those determined from TD-AMS data; LC/MS abundance was corrected by considering transmission efficiency.





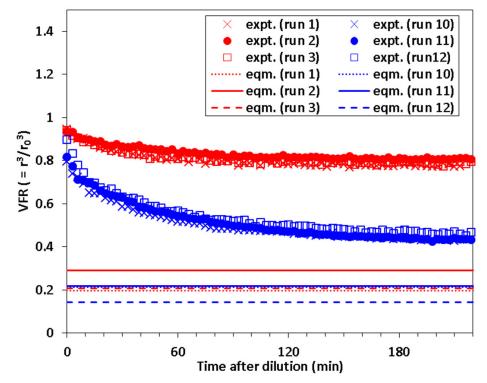


Figure 5: Time series of volume fraction remaining (VFR) measured for SOA particles after dilution under dry (RH < 1%; runs 1, 2, and 3) and humid conditions (RH  $\approx 40\%$ ; runs 10, 11, and 12); straight lines represent equilibrium levels calculated based on previous yield curve data.

470