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The maintenance of elevated active chlorine levels in the Antarctic lower stratosphere through HCl null-cycles

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Abstract. The Antarctic ozone hole arises from ozone destruction driven by elevated levels of ozone destroying ("active") chlorine in Antarctic spring. These elevated levels of active chlorine have to be formed first and then maintained throughout the period of ozone destruction. It is a matter of debate, how this maintenance of active chlorine is brought about in Antarctic spring, when the rate of formation of HCl (considered to be the main chlorine deactivation mechanism in Antarctica) is extremely high. Here we show that in the heart of the ozone hole (16-18 km or 100-70 hPa, in the core of the vortex), high levels of active chlorine are maintained by effective chemical cycles (referred to as HCl null-cycles hereafter). In these cycles, the formation of HCl is balanced by immediate reactivation, i.e. by immediate reformation of active chlorine. Under these conditions, polar stratospheric clouds sequester HNO_3 and thereby cause NO_2 concentrations to be low. These HCl null-cycles allow active chlorine levels to be maintained in the Antarctic lower stratosphere and thus rapid ozone destruction to occur. For the observed almost complete activation of stratospheric chlorine in the lower stratosphere, the heterogeneous reaction $HCl + HOCl$, the production of HOCl via $HO_2 + ClO$, with the HO_2 resulting from CH_2O photolysis, is essential. These results are important for assessing the impact of changes of the future stratospheric composition on the recovery of the ozone hole. Our simulations indicate that, in the lower stratosphere, future increased methane concentrations will not lead to enhanced chlorine deactivation (through the reaction $CH_4 + Cl \longrightarrow HCl + CH_3$) and that extreme ozone destruction to levels below ≈ 0.1 ppm will occur until mid-century.

1 Introduction

Because of the success of the Montreal protocol and its amendments and adjustments, the atmospheric halogen loading peaked in the early 1990s and has been declining slowly since then. However, even today, the Montreal Protocol has already achieved significant benefits for the ozone layer and the Antarctic ozone hole (by 2013, Chipperfield et al., 2015). Nonetheless, model simulations predict that the Antarctic ozone hole will continue to occur for decades (e.g., WMO, 2014; Oman et al., 2016; Fernandez et al., 2017).

The rapid chemical destruction of ozone in the Antarctic stratosphere in spring is caused by catalytic cycles driven by ClO and BrO (McElroy et al., 1986; Molina and Molina, 1987; Solomon, 1999). To run efficiently, these catalytic cycles require large concentrations of "active chlorine, ClO_x" ($ClO_x = ClO + 2 \times Cl_2O_2 + Cl$). The activation of chlorine, i.e., the conversion of the main chlorine reservoir species (HCl and $ClONO_2$) to ClO_x occurs through heterogeneous reactions (Solomon, 1999). The initial step of chlorine activation proceeds via the heterogeneous reaction (Solomon et al., 1986)

$$ClONO_2 + HCl \rightarrow Cl_2 + HNO_3 \quad ; \qquad\qquad\qquad\qquad\qquad\qquad\qquad\qquad (R1)$$

this step occurs very rapidly during polar night in about mid-May when temperatures become low enough for heterogeneous chlorine activation. Because of the initial concentration of HCl (before the onset of heterogeneous reactions) in the polar vortex being greater than that of $ClONO_2$ (Jaeglé et al., 1997; Santee et al., 2008), the amount of Cl_2 produced initially in reaction R1 is limited by the amount of available $ClONO_2$ (Salawitch et al., 1988; Crutzen et al., 1992; Portmann et al., 1996). The further chemical activation to near zero HCl values, as observed in Antarctic winter and in cold winters in the Arctic (Jaeglé et al., 1997; Santee et al., 2005; Santee et al., 2008; Manney et al., 2011; Wegner et al., 2012) requires the reproduction of partners (e.g. $ClONO_2$ or HOCl) for heterogeneous reactions with HCl.

With the return of sunlight to the polar region a period follows, characterised by further activation and maintenance of high levels of active chlorine (as observed, Santee et al., 2005; Santee et al., 2008) during which most of the ozone depletion occurs. Polar stratospheric clouds are measured in the Antarctic lower stratosphere until early October (Pitts et al., 2009). However, which chemical processes are responsible for further activation and maintenance of active chlorine is a matter of debate (Solomon et al., 2015).

Here, for the heart of the ozone layer (in the core of the vortex, in the lower stratosphere), where minimum ozone mixing ratios are reached (Solomon et al., 2005), we suggest the following picture of Antarctic ozone depletion. First, as the initial step of chlorine activation, the available $ClONO_2$ is titrated against HCl via Reaction R1. Then, very little chemical change occurs in polar night until early August ("sleeping chemistry") and a relatively slow additional chlorine activation until early September. The maintenance of high ClO_x values during mid-September to early October is accomplished by effective reaction cycles ("HCl null-cycles") which chemically inhibit a deactivation of chlorine that would otherwise proceed via net HCl formation. The period mid-September to early October is the period during which most of the Antarctic ozone loss occurs. This period of high ClO_x ends abruptly with the rapid formation of HCl leading to deactivation of chlorine (Grooß et al., 1997; Grooß et al., 2011).

We present box-model calculations of Antarctic chlorine chemistry and ozone depletion which allow the chemical mechanisms and the impact of particular reactions to be studied in detail. In particular, we apply a detailed analysis of the temporal development of the rates of the key chemical processes based on a unique algorithm for the determination of chemical pathways (Lehmann, 2004). We demonstrate that for the efficacy of the HCl null-cycles it is essential that the heterogeneous reaction (Prather, 1992; Crutzen et al., 1992)

$$HOCl + HCl \rightarrow Cl_2 + H_2O \qquad\qquad\qquad\qquad\qquad\qquad\qquad\qquad\qquad (R2)$$

as well as the gas-phase reaction (Crutzen et al., 1992)

$$CH_3O_2 + ClO \rightarrow CH_3O + ClOO \tag{R3}$$

(where the product ClOO decomposes rapidly to Cl and O_2) occurs. Further, the related formation of HO_x radicals from CH_3O photolysis is important.

2 Methods

2.1 Model description

The simulations presented here were performed with the Chemical Lagrangian model of the Stratosphere (McKenna et al., 2002; Grooß et al., 2005) (CLaMS); the set-up follows closely one used earlier (Grooß et al., 2011; Zafar, 2016). The model is used in box-model mode, where stratospheric chemistry is calculated for air parcels along three-dimensional trajectories. The air parcels are defined by the location and time of minimum ozone soundings in the ozone hole period, from which trajectories are calculated both backward to June and forward to December. In this way, trajectories from the core of the vortex in the lower stratosphere are selected. The trajectories of the air parcels were calculated using wind and temperature data from operational analyses from the European Centre for Medium-range Weather Forecasts (ECMWF). The latitudinal range covered by the trajectory in June, July and August is roughly $60°S$ to $80°S$, in September roughly $70°S$ to $90°S$, and in October and November roughly $60°S$ to $85°S$. The diabatic descent rates were calculated using a radiation code (Morcrette, 1991; Zhong and Haigh, 1995) assuming a cloud-free atmosphere. We use temperatures from the ECMWF operational analyses and climatological ozone and water vapour profiles (Grooß and Russell, 2005).

The initial values for the main trace gases at the start of the simulation (1 June 2003) are: O_3 = 2.2 ppm, H_2O = 4.1 ppm, CH_4 = 1.2 ppm, HNO_3 = 4.5 ppb, HCl = 1.05 ppb, ClO_x = 1.01 ppb, $ClONO_2$ = 12 ppt, $HOCl$ = 5 ppt, Br_y = 17 ppt, CO = 16 ppb. The sensitivity of the results of the simulations on the initial ozone mixing ratio is discussed in the appendix. (All these values are given in molar mixing ratios; throughout this paper we use molar/volume mixing ratios).

In Antarctic winter, temperatures typically fall below the threshold for chlorine activation and for the occurrence of polar stratospheric clouds (PSCs) approximately in mid-May (Pitts et al., 2009). Consistently, the initial values for chlorine species assumed here imply that the initial titration of HCl and $ClONO_2$ (Reaction R1) has already occurred by 1 June. (The term titration is used here in the sense that in a reaction one of the reactants, here $ClONO_2$, is completely used up). Likewise, the initial value assumed for HNO_3 implies that denitrification through the sedimentation of large NAT (nitric acid trihydrate) particles (e.g., Fahey et al., 2001; Molleker et al., 2014; Grooß et al., 2014) had occurred by this time. This assumption likely constitutes a slightly too early onset of denitrification, but the impact of this assumption is minimal during polar night. The impact of denitrification has been explored in sensitivity studies (see appendix).

To integrate the system of stiff ordinary differential equations describing the chemistry we employ the solver SVODE (Brown et al., 1989) that does not use the family approximation. The chemical kinetic data are taken from Sander et al. (2011). The

photolysis rates are calculated in spherical geometry (Becker et al., 2000) for every hour using a climatological ozone profile for ozone hole conditions from HALOE measurements (Grooß and Russell, 2005).

Of particular importance for the simulations discussed here is the representation of the photolysis of CH_2O. We have employed the recommended setup for both cross sections and quantum yields for 223 K (Sander et al., 2011). Using the photolysis quantum yields suggested by Röth and Ehhalt (2015) produces very similar results to those presented here (Zafar, 2016).

The branching ratios for the two channels (molecular channel and radical channel) of the photolysis of CH_2O are uncertain, an estimate of the uncertainty, including possible line structures, is about $\pm$ 20% for the radical channel (Fig. 1 in Röth and Ehhalt, 2015). In a sensitivity study, we also employ branching ratios of 0 and 100% which cannot occur in reality (Röth and Ehhalt, 2015). They are used here as lower and upper limits. A possible temperature dependence of both the cross sections and the quantum yields of the photolysis of CH_2O could be important, as it could potentially lead to different production of HO_2 in the photolysis of CH_2O for the temperature range below 200 K relevant here. The intensities of the maxima of each absorption band increases with lower temperatures, but an accurate temperature dependence of these kinetic data for temperatures in the polar lower stratosphere cannot be considered here due to lack of laboratory information (Smith et al., 2006; Röth and Ehhalt, 2015).

Heterogeneous chemistry is calculated on ice, nitric acid trihydrate (NAT), liquid ternary particles ($H_2SO_4/H_2O/HNO_3$) and cold liquid binary aerosols. Temperature dependent uptake coefficients of heterogeneous reactions on liquid ternary and binary aerosols are taken from parametrisations by Shi et al. (2001) as recommended (Sander et al., 2011). For uptake coefficients for reactions on NAT the parametrisation of Carslaw and Peter (1997) is used, based on laboratory measurements of Hanson and Ravishankara (1993). NAT particles are assumed to form from supercooled ternary solutions (STS) droplets. A HNO_3 supersaturation of three (corresponding to about 10 K supercooling) is required for NAT formation. The NAT particle density is assumed to be $3 \cdot 10^{-3}$ cm^{-3}. This NAT particle density is lower than assumed by Grooß et al. (2011) leading to a greater impact of reactivity on liquid particles in the simulations presented here.

2.2 Pathway analysis

A pathway in a chemical reaction system is a set of reactions converting some reactants of interest into some products of interest through some intermediate species, for which no net production or destruction occurs. An integer factor ("multiplicity") may be assigned to each reaction. The algorithm used here for the automatic determination of all significant pathways in a chemical reaction system was developed by Lehmann (2004). As input it requires a set of chemical reaction equations and reaction rates, which are usually provided by a chemical model. Starting from the individual reactions (and their rates) as initial pathways, longer pathways are constructed step by step by connecting shorter ones. If a newly formed pathway contains sub-pathways, it is split into these. A rate for each pathway is calculated. Pathways with rates below a pre-described threshold are deleted already during the construction process, in order to avoid an intractably large number of pathways ("combinatorial explosion").

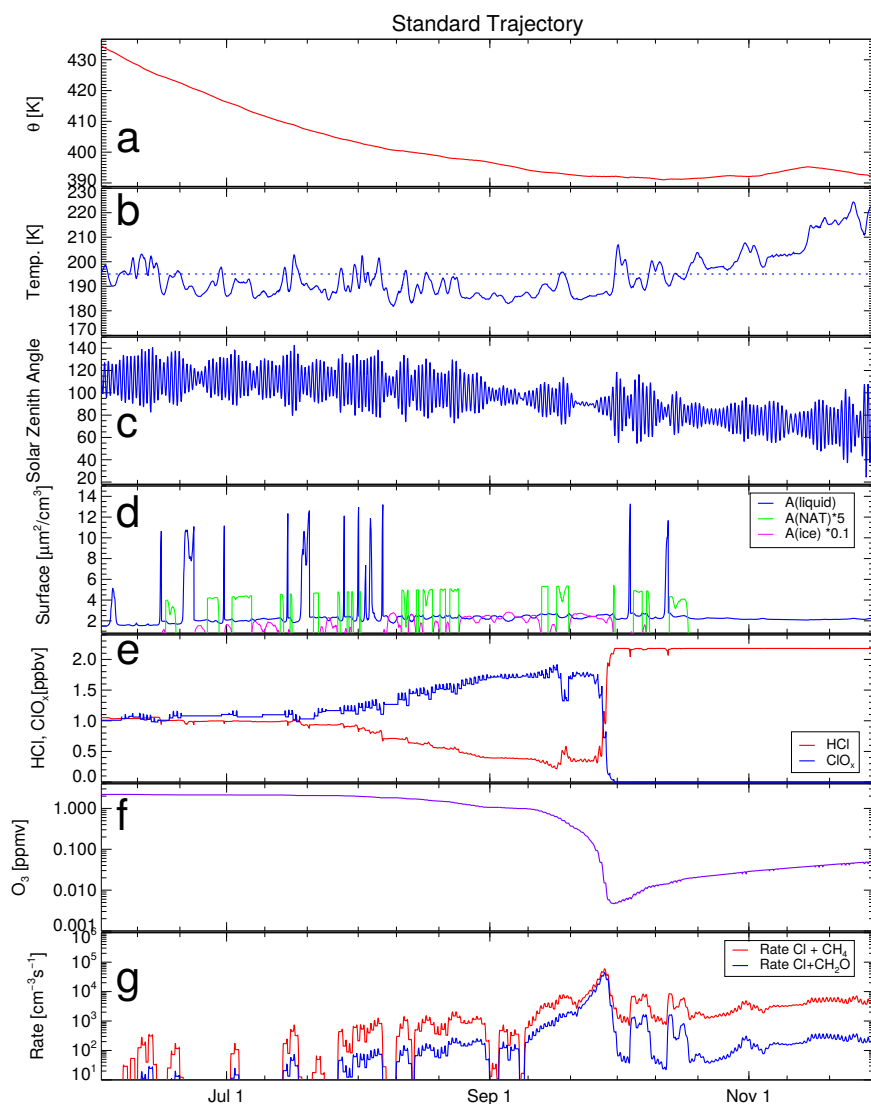



Figure 1. Box-model simulations along a trajectory passing through the location of the ozone sonde observation at South pole of 14 ppb on 74 hPa (391 K) on 24 September 2003 (Grooß et al., 2011). The different panels show a time series of the relevant parameters: (a) potential temperature of the air parcel, (b) temperature, (c) solar zenith angle, (d) surface area density of ice (red, scaled by 0.1), NAT (green, scaled by 5) and liquid aerosol particles (blue), (e) ClO_x (blue), and HCl (red), (f) ozone, and (g) rates of reaction of atomic chlorine with CH_4 (blue) and CH_2O (red). For reasons of clarity, the reaction rates in panel (g) are plotted as 24 h running averages. The model simulation covers the time period from 1 June to 30 November 2003.

3 Results

3.1 Maintenance of chlorine activation

We conducted box-model simulations for which the impact of mixing is neglected; however mixing across the Antarctic vortex edge is frequently overestimated substantially in current chemistry climate models (Hoppe et al., 2014). For the box-model, the development of temperature, potential temperature and solar zenith angle along a typical air parcel trajectory in Antarctic spring is shown in Figure 1 (panels a-c). The initial titration between HCl and $ClONO_2$ via Reaction R1 occurred before June (as usually in the Antarctic, e.g. Santee et al., 2008) so that the model simulation starts with about one ppb of ClO_x and with near zero values of $ClONO_2$. Little further chemical change occurs ("sleeping chemistry") as long as solar zenith angles are large (until end of July), but with decreasing solar zenith angle, HCl further decreases, leading to increasing ClO_x and, subsequently, chemical ozone destruction (e.g., Salawitch et al., 1988; Santee et al., 2005; Santee et al., 2008; Grooß et al., 2011). Ozone decreases most rapidly in September and reaches minimum values of less than 20 ppb on 24 September (for the specific case considered here).

With decreasing ozone, the Cl/ClO ratio shifts increasingly towards Cl (Douglass et al., 1995; Grooß et al., 1997), so that HCl production via the reactions of Cl with CH_4 and CH_2O (reactions R4 and R11) increases substantially by more than an order of magnitude between end of August and end of September (Figure 1, panel g). Indeed, HCl formation through reaction R4 is commonly considered as the main chlorine deactivation mechanism in Antarctic spring.

The formation rate of HCl via reactions R4 and R11 at the end of September amounts to more than 0.5 ppb per day. If these reactions proceeded unbalanced, chlorine activation and thus ozone depletion would be stopped within days. However, using an algorithm for the determination of chemical pathways (Lehmann, 2004) cycles C1 and C2 are identified for this period:

$$CH_4 + Cl \quad \rightarrow \quad HCl + CH_3 \tag{R4}$$

$$CH_3 + O_2 \quad \rightarrow \quad CH_3O_2 + M \tag{R5}$$

$$CH_3O_2 + ClO \quad \rightarrow \quad CH_3O + Cl + O_2 \tag{R3}$$

$$CH_3O + O_2 \quad \rightarrow \quad HO_2 + CH_2O \tag{R6}$$

$$ClO + HO_2 \quad \rightarrow \quad HOCl + O_2 \tag{R7}$$

$$HOCl + HCl \quad \rightarrow \quad Cl_2 + H_2O \tag{R2}$$

$$Cl_2 + h\nu \quad \rightarrow \quad 2\,Cl \tag{R8}$$

$$Cl + O_3 \quad \rightarrow \quad ClO + O_2 \quad (2\times) \tag{R9}$$

$$Net(C1): \quad CH_4 + 2\,O_3 \quad \rightarrow \quad CH_2O + H_2O + 2\,O_2 \tag{R10}$$

Cycle C1 was first formulated by Crutzen et al. (1992) demonstrating the importance of reaction R3 for ozone hole chemistry. Reaction R3 is essential for the production of a HO_2 radical from the CH_3 formed in reaction R4.

$$CH_2O + Cl \quad \rightarrow \quad HCl + CHO \tag{R11}$$

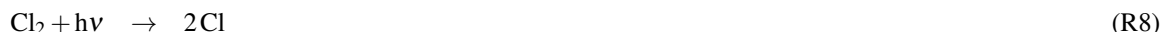
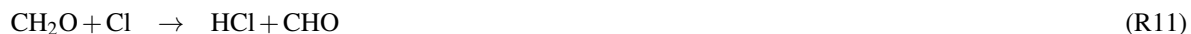
$$CHO + O_2 \quad \rightarrow \quad CO + HO_2 \tag{R12}$$

$$ClO + HO_2 \quad \rightarrow \quad HOCl + O_2 \tag{R7}$$

$$HOCl + HCl \quad \rightarrow \quad Cl_2 + H_2O \tag{R2}$$

$$Cl_2 + h\nu \quad \rightarrow \quad 2\,Cl \tag{R8}$$

$$Cl + O_3 \quad \rightarrow \quad ClO + O_2 \tag{R9}$$

$$Net(C2): \quad CH_2O + O_3 \quad \rightarrow \quad CO + H_2O + O_2 \tag{R13}$$

In cycles C1 and C2, HOCl is produced at the same rate as HCl because in both reactions R4 and R11, for each HCl molecule produced, also an HO_2 radical is generated (from CH_3 via R5, R3, and R6 or from CHO via R12). This HO_2 radical reacts further to form HOCl, which then reacts heterogeneously with HCl (in reaction R2) so that there is no net production of HCl. We therefore refer to cycles C1 and C2 in the following as "HCl null-cycles".

For the HCl null-cycle argument it is essential that Reaction R3 dominates the loss of CH_3O_2 and indeed alternative reactions for CH_3O_2 (e.g. the CH_3O_2 selfreaction) were discussed (Müller and Crutzen, 1994). However, all relevant alternative reactions are included in the chemical scheme used here and the competition of the different reaction pathways of CH_3O_2 is addressed in the pathway analysis that was employed (Lehmann, 2004).

These HCl null-cycles are effective in ensuring that there is no net production of HCl even though the speed of reactions R4 and R11, and thus the production of HCl, increases by about two orders of magnitude during September (Figure 1, panel g).

We find that during the first 10 days of September, HCl is recycled at a slow rate of 88 ppt in 10 days through cycle C1, at a rate of 26 ppt in 10 days through cycle C2. However, for the last days of September (21-30), the rate of recycling of HCl in cycles C1 and C2 is much more rapid, the rates are 3.7 ppb in 10 days and 2.9 ppb in 10 days respectively.

Thus cycles C1 and C2 constitute the chemical mechanism responsible for the maintenance of high levels of active chlorine (and thus of continued ozone destruction) under conditions of increasingly rapid HCl formation in reactions R4 and R11.

3.2 The path to full activation of HCl

As cycles C1 and C2 are HCl null-cycles, they cannot explain the decrease of HCl from values of ≈ 1 ppb in early July to very low values (about 0.2 ppb) during late August and September and thus the complete activation of chlorine as observed (Santee et al., 2008).

Using the pathway analysis (Lehmann, 2004), we identify two chemical cycles, C3 and C4 which are responsible for the decline of HCl during August and September:

$$CH_2O + h\nu \rightarrow CHO + H \tag{R14}$$

$$H + O_2 \rightarrow HO_2 + M \tag{R15}$$

$$CHO + O_2 \rightarrow CO + HO_2 \tag{R12}$$

$$ClO + HO_2 \rightarrow HOCl + O_2 \quad (2\times) \tag{R7}$$

$$HOCl + HCl \rightarrow Cl_2 + H_2O \quad (2\times) \tag{R2}$$

$$Cl_2 + h\nu \rightarrow 2\,Cl \quad (2\times) \tag{R8}$$

$$Cl + O_3 \rightarrow ClO + O_2 \quad (4\times) \tag{R9}$$

$$Net(C3): \quad CH_2O + 2\,HCl + 4\,O_3 \rightarrow CO + 2\,ClO + 2\,H_2O + 4\,O_2 \tag{R16}$$

$$O_3 + h\nu \rightarrow O(^1D) + O_2 \tag{R17}$$

$$O(^1D) + H_2O \rightarrow 2\,OH \tag{R18}$$

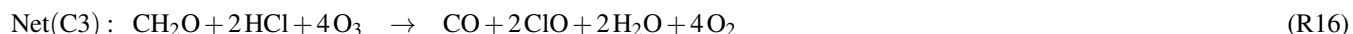
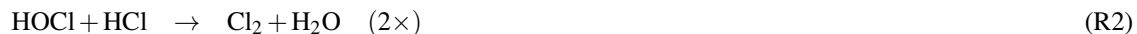
$$OH + O_3 \rightarrow HO_2 + O_2 \quad (2\times) \tag{R19}$$

$$ClO + HO_2 \rightarrow HOCl + O_2 \quad (2\times) \tag{R7}$$

$$HOCl + HCl \rightarrow Cl_2 + H_2O \quad (2\times) \tag{R2}$$

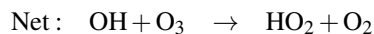
$$Cl_2 + h\nu \rightarrow 2\,Cl \quad (2\times) \tag{R8}$$

$$Cl + O_3 \rightarrow ClO + O_2 \quad (4\times) \tag{R9}$$

$$Net(C4): \quad 2\,HCl + 7\,O_3 \rightarrow 2\,ClO + H_2O + 9\,O_2 \tag{R20}$$

Reaction R19 may also proceed in two steps:

$$OH + ClO \rightarrow HO_2 + Cl \tag{R21}$$

$$Cl + O_3 \rightarrow ClO + O_2 \tag{R9}$$

$$Net: \quad OH + O_3 \rightarrow HO_2 + O_2$$

First, both cycles C3 and C4 require sufficiently fast heterogeneous reactions to be present. Second, in both cycles, it is important that HO_2 radicals are produced *without* simultaneous HCl formation (in contrast to cycles C1 and C2). The HO_2 radicals lead to formation of HOCl so that there is net HCl loss through reaction R2. The formation of HO_2 radicals in C3 and C4 is thus the key process responsible for the decline of HCl in August and September.

5   Reactions R9, R17, and R19 are fast enough to sustain cycles C3 and C4 even at the very low ozone concentrations as they occur in late September. For example, in the reference run for 26 September, the diurnal mean ozone mixing ratio is 77 ppb, the rate of reaction R9 is 46 ppb/d, the rate of reaction R17 is 15 ppb/d, and the net rate of reaction R19 is $3.2 \cdot 10^{-2}$ ppb/d

directly and 1.1 ppb/d when it proceeds in two steps (R21 and R9). A turnover of 46 ppb in reaction R9 in one day exceeds the available chlorine by more than an order of magnitude. This means that the available chlorine is cycled more than ten times per day through reaction R9. Therefore, even at extremely low ozone concentrations in late September, the rates of reactions R9, R17, and R19 are not rate limiting for cycles C3 and C4. Under these conditions, the rate limiting reaction for cycle C3 is the radical channel of the photolysis of CH_2O (R14) and for cycle C4 the reaction of $O(^1D)$ with water vapour (R18). (Note that reactions R9 and R17 do not constitute the rate limiting step of ozone loss cycles and can therefore not be used to deduce the ozone loss rate).

The dominant source of HO_x radicals under the conditions of the polar lower stratosphere in late winter and early spring is not the production of $O(^1D)$ radicals through ozone photolysis with subsequent reaction with H_2O, but rather the radical channel of the photolysis of CH_2O (reaction R14) (Crutzen et al., 1992; Müller and Crutzen, 1994; Crowley et al., 1994). Thus, the photolysis of CH_2O (radical channel) is effectively driving the depletion of HCl. In the time frame 20-31 August, 66 % of the net HCl depletion occurs through the photolysis of CH_2O and cycle C3 and only 15 % through the formation of $O(^1D)$ and subsequent reaction with H_2O (cycle C4). The rate of production of $ClONO_2$, under the conditions considered here, is very low (see discussion below).

The photolysis of CH_2O possesses two product channels; reaction R14 (radical channel) leading to the formation of two HO_2 radicals and the molecular channel

$$CH_2O + h\nu \rightarrow CO + H_2 \tag{R22}$$

which does not lead to production of HO_x. The branching ratio between reactions R14 and R22 is uncertain; it is about 30% for the radical channel R14 for the conditions in question here (see also section 2.1 Röth and Ehhalt, 2015).

To demonstrate the importance of cycle C3 and in particular the HO_2 production through reaction R14 and R15, we conducted two sensitivity runs, one assuming 100% efficiency for radical channel R14 and zero for the molecular channel R22 (red line in Figure 2) and vice versa (dark blue line in Figure 2). These assumptions constitute upper and lower limits of HO_2 production in reaction R14 and R15 that will certainly not be reached in reality.

Assuming no HO_2 production in the photolysis of CH_2O (blue line in Fig. 2) leads to very little reduction in HCl between late August and late September. This corroborates the conclusion that, in this chemical scheme, the HCl depletion beyond the initial titration against $ClONO_2$ is driven by the radical channel of the CH_2O photolysis (R14). Assuming 100% efficiency for the HO_2 producing channel in the CH_2O photolysis (red line in Figure 2) results in a much more rapid depletion of HCl than in the reference case and near zero values of HCl are reached in late August. Consequently, in this case, values of ClO_x are higher and ozone destruction is faster. Minimum values of ozone are reached somewhat earlier and likewise the corresponding rapid increase (deactivation) of HCl occurs somewhat earlier.

A more realistic estimate of the experimental uncertainties of the branching ratio for the radical channel of the CH_2O photolysis is $\pm 20\%$ (see section 2.1). Results for this estimate are also shown in Fig. 2, light blue and yellow lines. The results of this sensitivity test show enhanced HCl depletion for an enhanced radical channel (and reduced HCl depletion for a reduced radical channel), albeit to a much lower extent than for the upper and lower limit assumption.

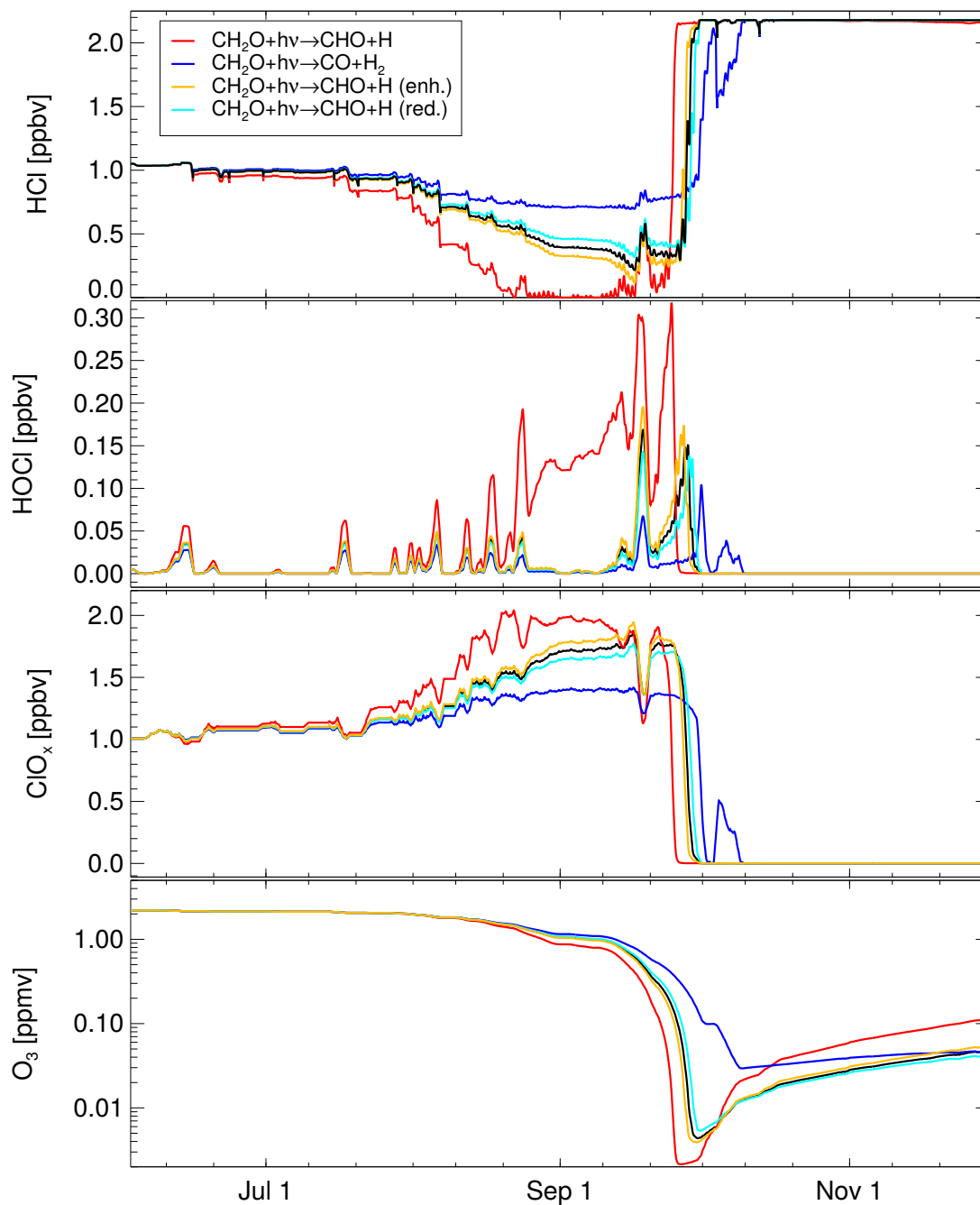



Figure 2. Sensitivity of the ozone hole chemistry on HO_2 production in the photolysis of CH_2O. Black line shows the reference case, red line a case assuming 100% efficiency for HO_2 production (i.e. for reaction R14), and dark blue line a case assuming no HO_2 production in the photolysis of CH_2O (100% efficiency for reaction R22). The light blue and yellow line show a more realistic estimate of the kinetic uncertainties, namely branching ratio for the radical channel enhanced by 20% (yellow line) and reduced by 20% (light blue line).

During the period of near zero values of HCl during late August to mid September (in the case of an 100 % radical channel) no substantial decrease of ClO_x and thus no deactivation through an increase in HCl occurs (red line in Figure 2). This is the case because cycles C1 and C2 (net reactions R10 and R13) are still active and prevent a net formation of HCl. Solely the ongoing production of HO_2 radicals causes somewhat enhanced levels of HOCl (0.1 to 0.2 ppb) during this period. Photolysis of HOCl inhibits an accumulation of larger amounts of HOCl.

3.3 The role of $ClONO_2$ and denitrification

For the conditions in the heart of the ozone layer, in the lower stratosphere, which are considered here, reaction

$$ClO + NO_2 + M \rightarrow ClONO_2 + M \tag{R23}$$

and the related chemistry involving $ClONO_2$ only play a minor role. This is the case mainly because PSCs consisting of NAT or STS exist almost continuously throughout the simulation sequestering most HNO_3 from the gas-phase. This leads to low NO_2 concentrations even in the presence of sunlight in spring. Consequently, throughout Antarctic winter and spring, and in particular during the main ozone loss period from late August to late September, the rate of the heterogeneous reaction between HCl and HOCl (R2) is substantially larger than the rate of HCl + $ClONO_2$ (R1), which is caused by the slow rate of formation of $ClONO_2$. The ClO radical preferentially reacts with HO_2 (forming HOCl) rather than with NO_2 (forming $ClONO_2$).

For the efficacy of cycles C1 and C2, low gas-phase concentrations of NO_2 (and thus of HNO_3) are necessary. It is however not important whether the removal of HNO_3 from the gas-phase occurs temporarily (uptake in particles) or permanently (denitrification). In our box-model simulations the impact of denitrification is taken into account by assuming that denitrification has occurred by the start of the simulation (4.5 ppb initial HNO_3 at 1 June). Overall, for the conditions in the core of the Antarctic vortex considered here, there is very little impact of denitrification on ozone depletion (see appendix).

3.4 Sensitivity of ozone loss on stratospheric methane and chlorine levels

In the coming decades, the composition of the Antarctic lower stratosphere will change considerably, in particular the stratospheric halogen loading will continue to decrease (e.g., WMO, 2014; Chipperfield et al., 2015). Repeating the reference run with Cl_y halved (blue line in Figure 3) (typical conditions for $\approx$2050) results in lower active chlorine ClO_x and thus slower ozone loss rates, as expected. However, in this case the ozone depletion period in the lower stratosphere is longer, so that extremely low ozone values (below 0.2 ppb) are reached also for the initial amount of Cl_y halved, albeit about three weeks later than in the reference case. The fact that in our simulation the impact of a reduced amount of Cl_y is mainly observed in September is consistent with the recent conclusion that signs of healing in the Antarctic ozone layer have emerged for September (Solomon et al., 2016). Note that in the reference case ozone depletion stops because ozone is reaching extremely low values although temperatures are still low enough for PSCs and heterogeneous processing; the same mechanism works for halved Cl_y, solely the low values are reached later in the season due to the slower ozone loss rate.

Doubling methane in the simulation leads to a speed up of Reaction R4, thereby also to enhanced production of CH_2O and thus also to a speed up of Reaction R11. But doubling methane does not lead to an enhanced deactivation, because cycles

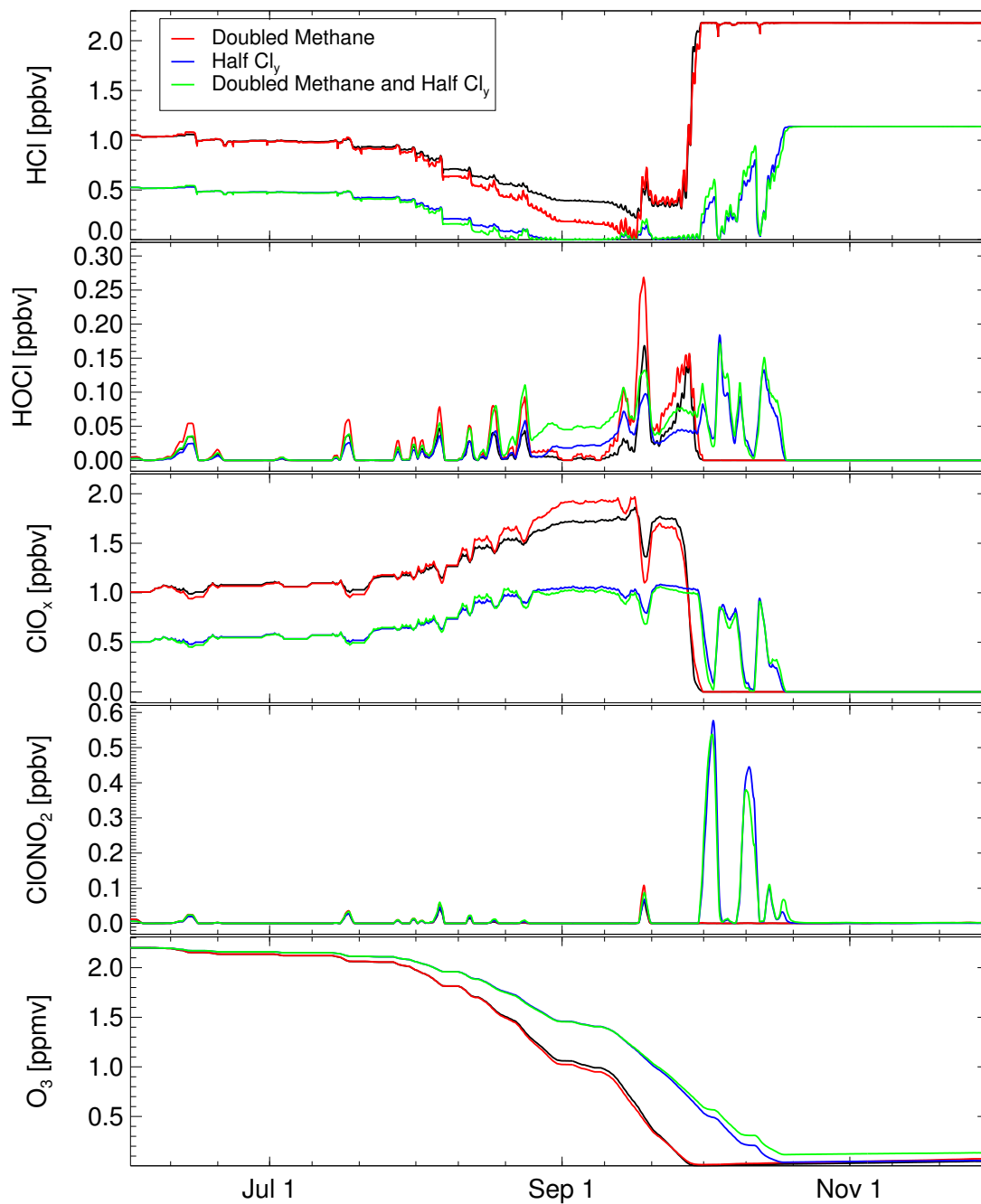



Figure 3. Sensitivity of the ozone hole chemistry on stratospheric methane and chlorine levels. Black line shows the reference case, red line a case assuming initial methane to be doubled, blue line a case assuming initial inorganic chlorine (Cl_y) to be halved, and green line a case for initial methane doubled and Cl_y halved.

C1 and C2 are active and inhibit the deactivation effect of Reactions R4 and R11. However, surprisingly at first, doubling of methane means that more CH_2O is produced in cycle C1 leading to somewhat faster HCl activation in late August and early September through cycle C3 and thus to higher ClO_x and a somewhat faster ozone depletion during this period.

A simulation combining doubled methane and halved Cl_y shows results very close to those for halved Cl_y, but again the doubled methane does not lead to a substantially faster deactivation and practically the same very low ozone values are reached in late October. Thus we suggest that the very low ozone values observed today in the core of the polar vortex (Solomon et al., 2005) will continue to occur for decades to come and that the recovery of the stratospheric chlorine loading should lead to a shift by some weeks of the low values to later in the season, consistent with recent observations (Solomon et al., 2016).

A feature only noticeable in the simulations for halved Cl_y are peaks of enhanced $ClONO_2$ in late September and early October (Fig. 3). At this time, for halved Cl_y, ozone is strongly reduced, but clearly higher than in the reference run and full deactivation has not occurred. During this time period, PSCs occasionally evaporate (owing to higher temperatures, Fig. 1), releasing HNO_3 to the gas-phase. This enhanced HNO_3 leads to the production of NO_2 with subsequent formation of $ClONO_2$, similar to the formation of $ClONO_2$ in the Arctic in late winter and early spring in periods with higher temperatures when PSCs evaporate (e.g., Müller et al., 1994). However, as discussed above, for halved Cl_y extremely low ozone mixing ratios are reached (albeit later than in the reference case), so that the final chlorine deactivation in mid October occurs through a rapid, practically complete conversion of active chlorine into HCl (Fig. 3).

3.5 Multi-trajectory simulations

We have repeated the reference run using the CLaMS model (McKenna et al., 2002; Grooß et al., 2005) in box-model mode for multiple trajectories. We employ a set of realistic trajectories passing the South pole at 400 K potential temperature (in late September/early October) including diabatic descent and latitude variations (taken from Grooß et al., 2011). In the period early August to early October all trajectories are subject to roughly the same diabatic descent of ≈ 10 K, similarly as for the reference run. In this period, all trajectories show strong variations in latitude, again similar as for the reference run. The latitude varies between the South Pole and $\approx 65°$S with some equatorward excursions to $\approx 60°$S or, even more rarely, to $\approx 55°$S.

The initial values (for 1 August) of HCl and ClO_x were chosen consistently with the reference run. The results of these runs (Fig. 4) show a significant variability in the decline of HCl and ozone loss rate, but all show a rather similar behaviour as the reference run (shown in red). Namely, very little chemical change in polar night until early August ("sleeping chemistry") and a relatively slow additional chlorine activation until early September. Next, a maintenance of high ClO_x values during mid-September to early October (due to cycles C1 and C2) accompanied by rapid ozone loss. This period of high ClO_x ends abruptly with the rapid formation of HCl (and thus chlorine deactivation), which occurs in a similar way for the individual trajectories in the time frame late September to early October.

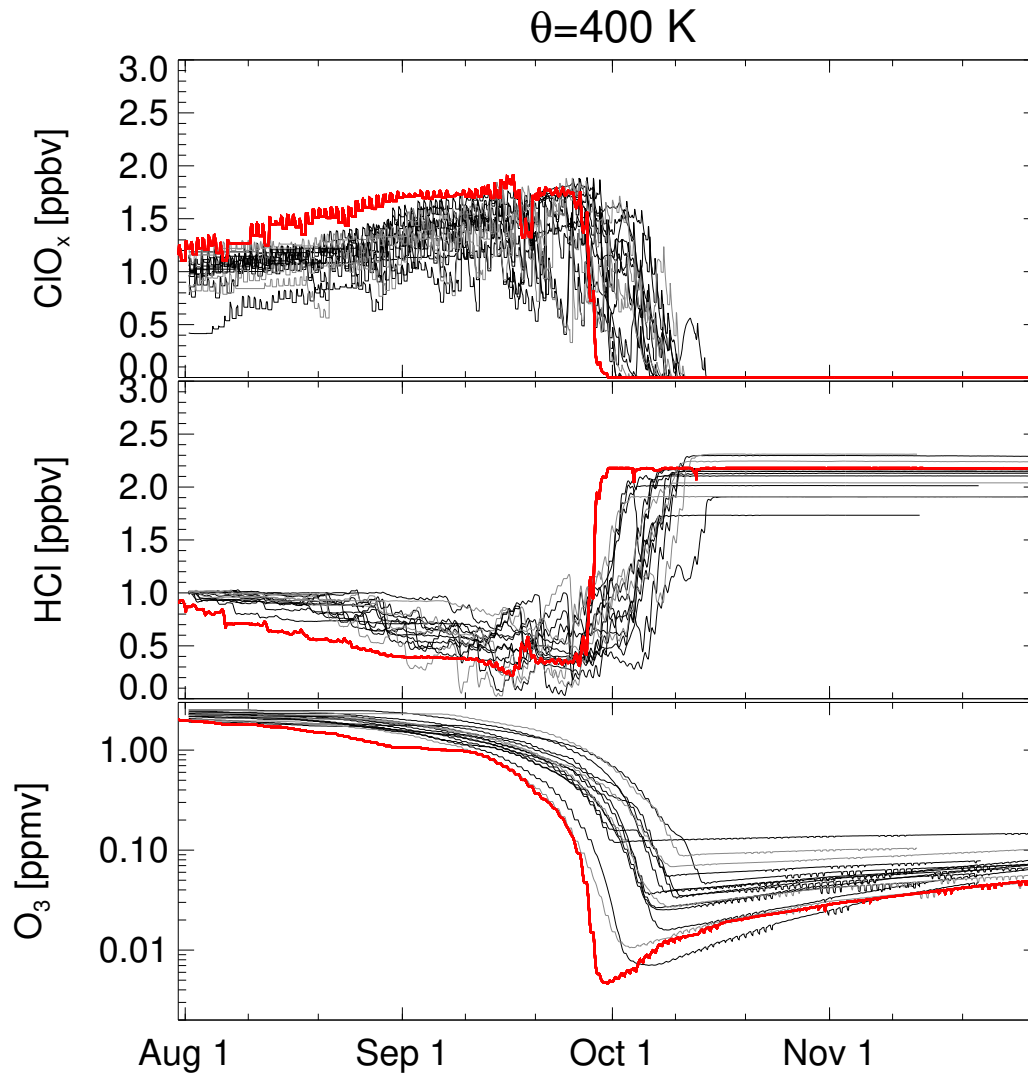



Figure 4. Results from multi-trajectory simulations of CLaMS. Box-model simulations were performed for a set of trajectories passing the South pole at 400 K taken from Grooß et al. (2011). The box-model simulations cover the time period from 1 August to 30 November 2003. A few trajectories showing very little diabatic descent (and thus much smaller values of total chlorine) were neglected. Initial values were chosen consistently with the reference run (Fig. 1), which is shown here in red. Individual trajectories are shown in different shades of grey to allow them to be distinguished more easily.

4 Conclusions

For the heart of the ozone hole in the Antarctic lowermost stratosphere, formation of HCl through reactions R4 and R11 is very rapid for enhanced chlorine levels with the rate of HCl formation increasing by more than a factor of ten during September. We have shown that high levels of active chlorine are maintained nonetheless, because the formation of HCl is balanced in very effective HCl null-cycles, allowing rapid chemical destruction of ozone to proceed. Further, for the depletion of HCl to very low values, formation of HO_2 is essential, with photolysis of CH_2O being the major net source of HO_2. Owing to the uptake of HNO_3 in NAT and STS particles, NO_x chemistry and the formation of $ClONO_2$ are of minor importance. Chlorine is finally deactivated if an imbalance in the null-cycles occurs leading to a rapid and almost complete conversion of the activated chlorine into HCl and thereby putting a halt to ozone depletion (Douglass et al., 1995; Grooß et al., 1997; Grooß et al., 2011). An increase of methane in the future should not lead to a faster ozone recovery as might be expected. These results are important for an assessment of the impact of chemical change to come in the Antarctic stratosphere on the future development of the ozone hole.

Data availability. The model results presented here can be obtained in electronic form (netcdf-files) from the corresponding author on request.

Author contributions. R.M., J.-U.G., and R.L. conceived and designed the research project. A.M.Z., S.R., and J.-U.G. conducted the simulations. R.L. performed the pathway analysis. R.M., J.-U.G., A.M.Z., S.R., and R.L. contributed to the interpretation of the results and wrote the paper

Competing interests. The authors declare that they have no competing financial interests

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Appendix A: Sensitivity studies

In this section we discuss the sensitivity of the simulated development of chlorine chemistry and ozone mixing ratios with respect to the initial ozone and HNO$_3$ mixing ratios (similar as in Grooß et al., 2011).

A1   Initial ozone mixing ratio
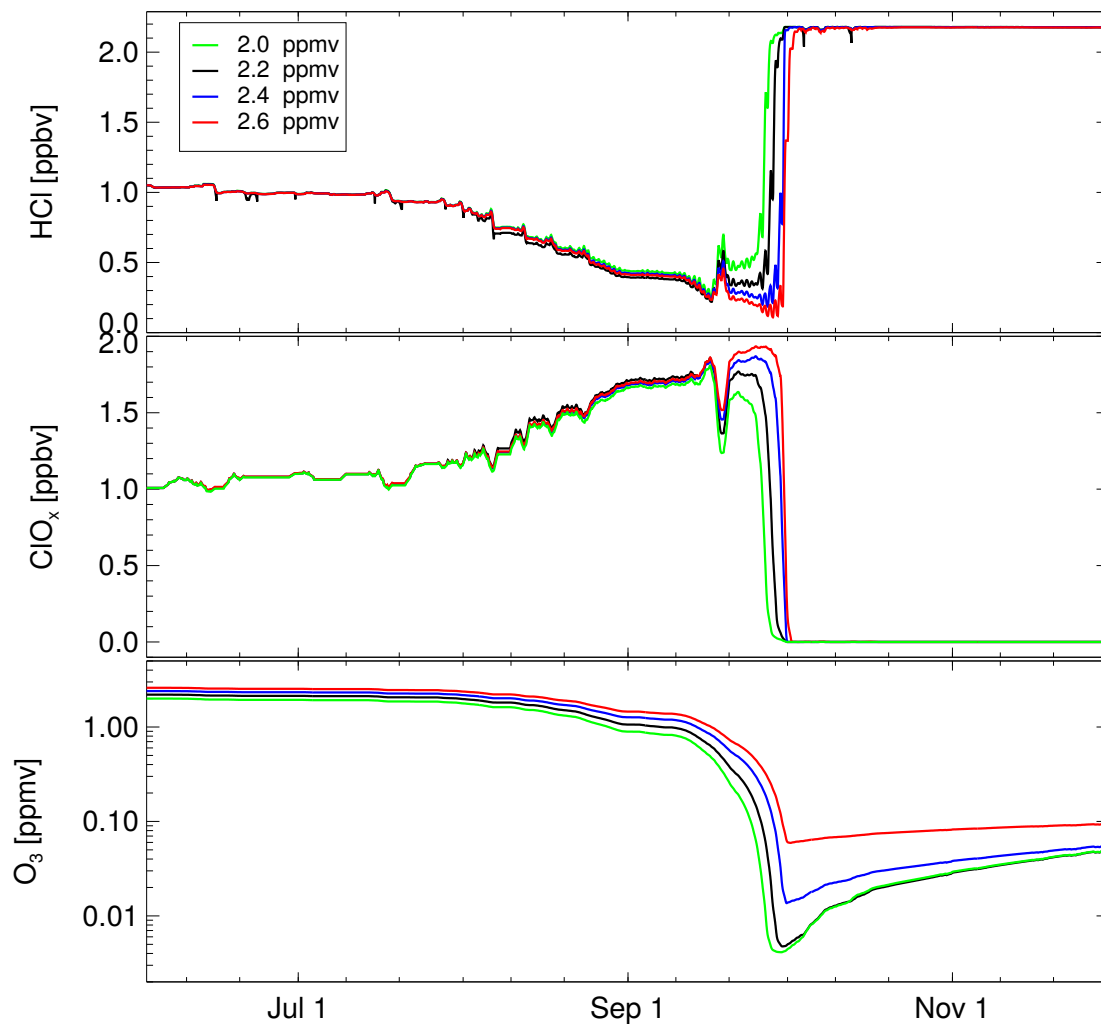
Figure A1. Simulations for different initial ozone mixing ratios. Simulation as in the reference run but with different initial ozone mixing ratios (2.0 to 2.6 ppm).

5   We conducted three additional box-model simulations, identical to the reference run (initial ozone 2.2 ppm), but with different initial ozone mixing ratios ranging from 2.0 to 2.6 ppm (Fig. A1). The simulations show both different minimum ozone mixing ratios and different times at which these minima are reached (consistent with the results of Grooß et al., 2011). Simi-

larly, the timing of the rapid increase in HCl and chlorine deactivation is different for the different initial ozone mixing ratios (Fig. A1).

Importantly however, the general pattern of chlorine activation and decrease of HCl is very similar until about 20 September indicating that the efficacy of the HCl null-cycles is not affected by the initial ozone mixing ratio.

A2   Denitrification
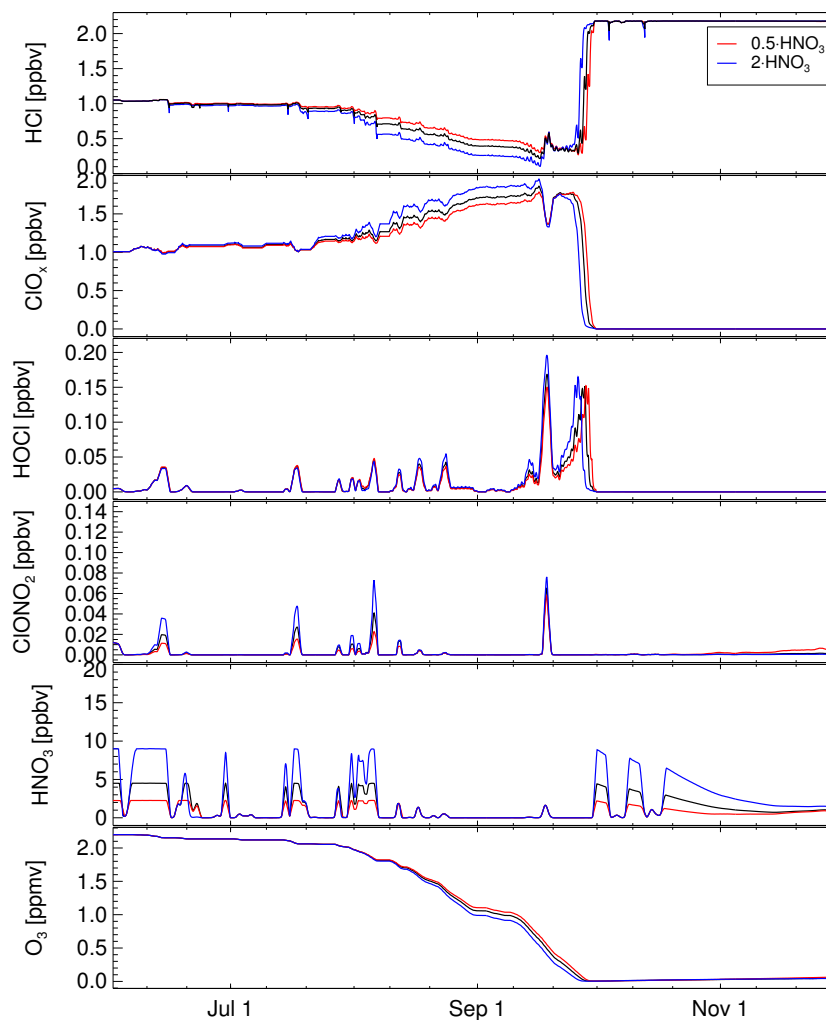
Figure A2. Simulation for the impact of denitrification. The impact of denitrification is tested by initialising the model run with HNO_3 doubled (blue line) and halved (red line). Black line shows the results of the reference simulation.

The formation of PSCs impacts polar chemistry by taking up HNO_3 and thereby reducing HNO_3 in the gas-phase (Crutzen and Arnold, 1986; Toon et al., 1986; Salawitch et al., 1988). In the case of STS particles the reduction of gas-phase HNO_3

lasts only as long as temperatures are low enough for the particles to exist, because the particles are too small to sediment significantly. Nitric acid trihydrate particles, however, can grow to large sizes allowing substantial sedimentation rates, so that HNO_3 is removed from stratospheric air masses permanently (a process referred to as "denitrification", Fahey et al., 2001; Molleker et al., 2014). Denitrification is observed regularly in the lower stratosphere in Antarctica and in cold winters in the

5 Arctic (e.g., Davies et al., 2006; Santee et al., 2008; Grooß et al., 2014).

For the case considered here, PSCs that sequester HNO_3 from the gas-phase exist almost continuously until mid September (Fig. 1). Therefore, the gas-phase concentration of HNO_3 during the period of chlorine activation, maintenance of active chlorine, and chemical ozone depletion is controlled by PSCs and not by denitrification.

Nonetheless, we conducted a sensitivity simulation initialised with half (2.25 ppb) or twice (9 ppb) the amount of initial

10 HNO_3 (Fig. A2). These assumptions result in practically the same ozone depletion in late September as in the reference case (Fig. A2). For mid September, the simulation with 9 ppb initial HNO_3 shows about 10% more active chlorine than the reference run and thus a somewhat more rapid ozone loss (ozone in mid September ≈ 0.1 ppm less than in the reference run). For half the initial HNO_3 (2.25 ppb) the situation is reversed, but differences are even smaller.

Thus, as to be expected, the impact of the initial HNO_3 mixing ratio on the simulated ozone loss and chlorine activation

15 is only moderate. Solely in the short periods when HNO_3 is released to the gas-phase and some of this gas-phase HNO_3 is converted to NO_x, the resulting production of $ClONO_2$ is enhanced in the case of a greater initial HNO_3 mixing ratio. Therefore, a greater initial HNO_3 mixing ratio (assumption of no denitrification) leads to a somewhat faster HCl depletion (through R1), a stronger chlorine activation and thus a little bit more chemical ozone loss (Fig. A2).

Overall, like for the sensitivity on the initial ozone mixing ratio, the general pattern of chlorine activation and decrease of

20 HCl is very similar for the results of the reference run and the results under the (rather extreme) assumptions of reduced and enhanced denitrification (Fig. A2).

References

Becker, G., Grooß, J.-U., McKenna, D. S., and Müller, R.: Stratospheric photolysis frequencies: Impact of an improved numerical solution of the radiative transfer equation, J. Atmos. Chem., 37, 217–229, https://doi.org/10.1023/A:1006468926530, 2000.

Brown, P. N., Byrne, G. D., and Hindmarsh, A. C.: VODE: A variable coefficient ODE solver, SIAM J. Sci. Stat. Comput., 10, 1038–1051, 1989.

Carslaw, K. S. and Peter, T.: Uncertainties in reactive uptake coefficients for solid stratospheric particles - 1. Surface chemistry, Geophys. Res. Lett., 24, 1743–1746, 1997.

Chipperfield, M. P., Dhomse, S. S., Feng, W., McKenzie, R. L., Velders, G. J. M., and Pyle, J. A.: Quantifying the ozone and ultraviolet benefits already achieved by the Montreal Protocol, Nat. Commun., 6, 7233, https://doi.org/10.1038/ncomms8233, 2015.

Crowley, J. N., Helleis, F., Müller, R., Moortgat, G. K., Crutzen, P. J., and Orlando, J. J.: CH_3OCl: UV/Vis absorption cross-sections, J values and atmospheric significance, J. Geophys. Res., 99, 20 683–20 688, 1994.

Crutzen, P. J. and Arnold, F.: Nitric acid cloud formation in the cold Antarctic stratosphere: A major cause for the springtime 'ozone hole', Nature, 342, 651–655, 1986.

Crutzen, P. J., Müller, R., Brühl, C., and Peter, T.: On the potential importance of the gas phase reaction $CH_3O_2 + ClO \rightarrow ClOO + CH_3O$ and the heterogeneous reaction $HOCl + HCl \rightarrow H_2O + Cl_2$ in "ozone hole" chemistry, Geophys. Res. Lett., 19, 1113–1116, https://doi.org/10.1029/92GL01172, 1992.

Davies, S., Mann, G. W., Carslaw, K. S., Chipperfield, M. P., Remedios, J. J., Allen, G., Waterfall, A. M., Spang, R., and Toon, G. C.: Testing our understanding of Arctic denitrification using MIPAS-E satellite measurements in winter 2002/2003, Atmos. Chem. Phys., 6, 3149–3161, https://doi.org/10.5194/acp-6-3149-2006, 2006.

Douglass, A. R., Schoeberl, M. R., Stolarski, R. S., Waters, J. W., Russell III, J. M., Roche, A. E., and Massie, S. T.: Interhemispheric differences in springtime production of HCl and $ClONO_2$ in the polar vortices, J. Geophys. Res., 100, 13 967–13 978, 1995.

Fahey, D. W., Gao, R. S., Carslaw, K. S., Kettleborough, J., Popp, P. J., Northway, M. J., Holecek, J. C., Ciciora, S. C., McLaughlin, R. J., Thompson, T. L., Winkler, R. H., Baumgardner, D. G., Gandrud, B., Wennberg, P. O., Dhaniyala, S., McKinley, K., Peter, T., Salawitch, R. J., Bui, T. P., Elkins, J. W., Webster, C. R., Atlas, E. L., Jost, H., Wilson, J. C., Herman, R. L., Kleinböhl, A., and von König, M.: The detection of large HNO_3-containing particles in the winter Arctic stratosphere, Science, 291, 1026–1031, 2001.

Fernandez, R. P., Kinnison, D. E., Lamarque, J.-F., Tilmes, S., and Saiz-Lopez, A.: Impact of biogenic very short-lived bromine on the Antarctic ozone hole during the 21st century, Atmos. Chem. Phys., 17, 1673–1688, https://doi.org/10.5194/acp-17-1673-2017, http://www.atmos-chem-phys.net/17/1673/2017/, 2017.

Grooß, J.-U. and Russell, J. M.: Technical note: A stratospheric climatology for O_3, H_2O, CH_4, NO_x, HCl, and HF derived from HALOE measurements, Atmos. Chem. Phys., 5, 2797–2807, 2005.

Grooß, J.-U., Pierce, R. B., Crutzen, P. J., Grose, W. L., and Russell III, J. M.: Re-formation of chlorine reservoirs in southern hemisphere polar spring, J. Geophys. Res., 102, 13 141–13 152, https://doi.org/10.1029/96JD03505, 1997.

Grooß, J.-U., Günther, G., Müller, R., Konopka, P., Bausch, S., Schlager, H., Voigt, C., Volk, C. M., and Toon, G. C.: Simulation of denitrification and ozone loss for the Arctic winter 2002/2003, Atmos. Chem. Phys., 5, 1437–1448, 2005.

Grooß, J.-U., Brautzsch, K., Pommrich, R., Solomon, S., and Müller, R.: Stratospheric ozone chemistry in the Antarctic: What controls the lowest values that can be reached and their recovery?, Atmos. Chem. Phys., 11, 12 217–12 226, 2011.

Grooß, J.-U., Engel, I., Borrmann, S., Frey, W., Günther, G., Hoyle, C. R., Kivi, R., Luo, B. P., Molleker, S., Peter, T., Pitts, M. C., Schlager, H., Stiller, G., Vömel, H., Walker, K. A., and Müller, R.: Nitric acid trihydrate nucleation and denitrification in the Arctic stratosphere, Atmos. Chem. Phys., 14, 1055–1073, https://doi.org/10.5194/acp-14-1055-2014, 2014.

Hanson, D. R. and Ravishankara, A. R.: Reaction of $ClONO_2$ with HCl on NAT, NAD, and frozen sulfuric acid and hydrolysis of N_2O_5 and $ClONO_2$ on frozen sulfuric acid, J. Geophys. Res., 98, 22 931–22 936, 1993.

Hoppe, C. M., Hoffmann, L., Konopka, P., Grooß, J.-U., Ploeger, F., Günther, G., Jöckel, P., and Müller, R.: The implementation of the CLaMS Lagrangian transport core into the chemistry climate model EMAC 2.40.1: application on age of air and transport of long-lived trace species, Geosci. Model Dev., 7, 2639–2651, https://doi.org/10.5194/gmd-7-2639-2014, http://www.geosci-model-dev.net/7/2639/2014/, 2014.

Jaeglé, L., Webster, C. R., May, R. D., Scott, D. C., Stimpfle, R. M., Kohn, D. W., Wennberg, P. O., Hanisco, T. F., Cohen, R. C., Proffitt, M. H., Kelly, K. K., Elkins, J., Baumgardner, D., Dye, J. E., Wilson, J. C., Pueschel, R. F., Chan, K. R., Salawitch, R. J., Tuck, A. F., Hovde, S. J., and Yung, Y. L.: Evolution and stoichiometry of heterogeneous processing in the Antarctic stratosphere, J. Geophys. Res., 102, 13 235–13 253, https://doi.org/10.1029/97JD00935, 1997.

Lehmann, R.: An algorithm for the determination of all significant pathways in chemical reaction systems, J. Atmos. Chem., 47, 45–78, 2004.

Manney, G. L., Santee, M. L., Rex, M., Livesey, N. J., Pitts, M. C., Veefkind, P., Nash, E. R., Wohltmann, I., Lehmann, R., Froidevaux, L., Poole, L. R., Schoeberl, M. R., Haffner, D. P., Davies, J., Dorokhov, V., Gernandt, H., Johnson, B., Kivi, R., Kyrö, E., Larsen, N., Levelt, P. F., Makshtas, A., McElroy, C. T., Nakajima, H., Parrondo, M. C., Tarasick, D. W., von der Gathen, P., Walker, K. A., and Zinoviev, N. S.: Unprecedented Arctic ozone loss in 2011, Nature, 478, 469–475, https://doi.org/10.1038/nature10556, 2011.

McElroy, M. B., Salawitch, R. J., Wofsy, S. C., and Logan, J. A.: Antarctic ozone: Reductions due to synergistic interactions of chlorine and bromine, Nature, 321, 759–762, 1986.

McKenna, D. S., Grooß, J.-U., Günther, G., Konopka, P., Müller, R., Carver, G., and Sasano, Y.: A new Chemical Lagrangian Model of the Stratosphere (CLaMS): 2. Formulation of chemistry scheme and initialization, J. Geophys. Res., 107, 4256, https://doi.org/10.1029/2000JD000113, 2002.

Molina, L. T. and Molina, M. J.: Production of Cl_2O_2 from the self-reaction of the ClO radical, J. Phys. Chem., 91, 433–436, 1987.

Molleker, S., Borrmann, S., Schlager, H., Luo, B., Frey, W., Klingebiel, M., Weigel, R., Ebert, M., Mitev, V., Matthey, R., Woiwode, W., Oelhaf, H., Dörnbrack, A., Stratmann, G., Grooß, J.-U., Günther, G., Vogel, B., Müller, R., Krämer, M., Meyer, J., and Cairo, F.: Micro-physical properties of synoptic-scale polar stratospheric clouds: in situ measurements of unexpectedly large HNO_3-containing particles in the Arctic vortex, Atmos. Chem. Phys., 14, 10 785–10 801, https://doi.org/10.5194/acp-14-10785-2014, http://www.atmos-chem-phys.net/14/10785/2014/, 2014.

Morcrette, J.-J.: Radiation and cloud radiative properties in the European Centre for Medium-Range Weather Forecasts forecasting system, J. Geophys. Res., 96, 9121–9132, 1991.

Müller, R. and Crutzen, P. J.: On the relevance of the methane oxidation cycle to "ozone hole" chemistry, in: Ozone in the troposphere and stratosphere, edited by Hudson, R. D., Proceedings of the Quadrennial Ozone Symposium 1992, pp. 298–301, 1994.

Müller, R., Peter, T., Crutzen, P. J., Oelhaf, H., Adrian, G. P., v. Clarmann, T., Wegner, A., Schmidt, U., and Lary, D.: Chlorine chemistry and the potential for ozone depletion in the Arctic stratosphere in the winter of 1991/92, Geophys. Res. Lett., 21, 1427–1430, 1994.

Oman, L. D., Douglass, A. R., Salawitch, R. J., Canty, T. P., Ziemke, J. R., and Manyin, M.: The effect of representing bromine from VSLS on the simulation and evolution of Antarctic ozone, Geophys. Res. Lett., 43, 9869–9876, https://doi.org/10.1002/2016GL070471, 2016.

Pitts, M. C., Poole, L. R., and Thomason, L. W.: CALIPSO polar stratospheric cloud observations: second-generation detection algorithm and composition discrimination, Atmos. Chem. Phys., 9, 7577–7589, 2009.

Portmann, R. W., Solomon, S., Garcia, R. R., Thomason, L. W., Poole, L. R., and McCormick, M. P.: Role of aerosol variations in anthropogenic ozone depletion in the polar regions, J. Geophys. Res., 101, 22 991–23 006, 1996.

Prather, M. J.: More rapid ozone depletion through the reaction of HOCl with HCl on polar stratospheric clouds, Nature, 355, 534–537, 1992.

Röth, E.-P. and Ehhalt, D. H.: A simple formulation of the CH_2O photolysis quantum yields, Atmos. Chem. Phys., 15, 7195–7202, https://doi.org/10.5194/acp-15-7195-2015, http://www.atmos-chem-phys.net/15/7195/2015/, 2015.

Salawitch, R. J., Wofsy, S. C., and McElroy, M. B.: Influence of polar stratospheric clouds on the depletion of Antarctic ozone, Geophys. Res. Lett., 15, 871–874, 1988.

Sander, S. P., Abbatt, J., Barker, J. R., Burkholder, J. B., Friedl, R. R., Golden, D. M., Huie, R. E., Kolb, C. E., Kurylo, M. J., Moortgat, G. K., Orkin, V. L., and Wine, P. H.: Evaluation No. 17, JPL Publication 10-6, Chemical Kinetics and Photochemical Data for Use in Atmospheric Studies, Jet Propulsion Laboratory, Pasadena, http://jpldataeval.jpl.nasa.gov, 2011.

Santee, M. L., Manney, G. L., Livesey, N. J., Foidevaux, L., MacKenzie, I. A., Pumphrey, H. C., Read, W. G., Schwartz, M. J., Waters, J. W., and Harwood, R. S.: Polar processing and development of the 2004 Antarctic ozone hole: First results from MLS on Aura, Geophys. Res. Lett., 32, 2005.

Santee, M. L., MacKenzie, I. A., Manney, G. L., Chipperfield, M. P., Bernath, P. F., Walker, K. A., Boone, C. D., Froidevaux, L., Livesey, N. J., and Waters, J. W.: A study of stratospheric chlorine partitioning based on new satellite measurements and modeling, J. Geophys. Res., 113, D12307, https://doi.org/10.1029/2007JD009057, 2008.

Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: Kinetic model for reaction of $ClONO_2$ with H_2O and HCl and HOCl with HCl in sulfuric acid solutions, J. Geophys. Res., 106, 24 259–24 274, https://doi.org/10.1029/2000JD000181, 2001.

Smith, C. A., Pope, F. D., Cronin, B., Parkes, C. B., and Orr-Ewing, A. J.: Absorption cross sections of formaldehyde at wavelengths from 300 to 340 nm at 294 and 245 K, J. Phys. Chem. A, 110, 11 645–11 653, https://doi.org/10.1021/jp063713y, pMID: 17034158, 2006.

Solomon, S.: Stratospheric ozone depletion: A review of concepts and history, Rev. Geophys., 37, 275–316, https://doi.org/10.1029/1999RG900008, 1999.

Solomon, S., Garcia, R. R., Rowland, F. S., and Wuebbles, D. J.: On the depletion of Antarctic ozone, Nature, 321, 755–758, 1986.

Solomon, S., Portmann, R. W., Sasaki, T., Hofmann, D. J., and Thompson, D. W. J.: Four decades of ozonesonde measurements over Antarctica, J. Geophys. Res., 110, D21311, https://doi.org/10.1029/2005JD005917, 2005.

Solomon, S., Kinnison, D., Bandoro, J., and Garcia, R.: Simulation of polar ozone depletion: An update, J. Geophys. Res., 120, 7958–7974, https://doi.org/10.1002/2015JD023365, 2015.

Solomon, S., Ivy, D. J., Kinnison, D., Mills, M. J., Neely, R. R., and Schmidt, A.: Emergence of healing in the Antarctic ozone layer, Science, https://doi.org/10.1126/science.aae0061, 2016.

Toon, O. B., Hamill, P., Turco, R. P., and Pinto, J.: Condensation of HNO_3 and HCl in winter polar stratospheres, Geophys. Res. Lett., 13, 1284–1287, 1986.

Wegner, T., Grooß, J.-U., von Hobe, M., Stroh, F., Sumińska-Ebersoldt, O., Volk, C. M., Hösen, E., Mitev, V., Shur, G., and Müller, R.: Heterogeneous chlorine activation on stratospheric aerosols and clouds in the Arctic polar vortex, Atmos. Chem. Phys., 12, 11 095–11 106, https://doi.org/10.5194/acp-12-11095-2012, 2012.

WMO: Scientific assessment of ozone depletion: 2014, Global Ozone Research and Monitoring Project–Report No. 55, Geneva, Switzerland, 2014.

Zafar, A. M.: Stratospheric ozone depletion: Analysis of heterogeneous chemistry in the Antarctic, no. 4394 in Berichte des Forschungszentrums Jülich, Forschungszentrum Jülich, available from: www.fz-juelich.de/zb/openaccess, 2016.

5 Zhong, W. and Haigh, J. D.: Improved broadband emissivity parameterization for water vapor cooling rate calculations, J. Atmos. Sci., 52, 124–138, 1995.