



1 Quantification of atmospheric nucleation and growth process 2 as a single source of aerosol particles in a city

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6 **Abstract.** Effects of new aerosol particle formation (NPF) and particle diameter growth process as a single source on
7 atmospheric particle number concentrations were evaluated and quantified on the basis of experimental data sets obtained from
8 particle number size distribution measurements in the city centre and near-city background of Budapest for 5 years. Nucleation
9 strength factors separately for nucleation days ($NSF_{\text{nuc1 days}}$) and for all days ($NSF_{\text{all days}}$) were derived for seasons and full years.
10 The former characteristics represents the concentration increment of ultrafine (UF) particle numbers with respect to
11 background concentration due solely to nucleation specifically on nucleation days. The latter factor expresses the contribution
12 of nucleation process to the background particle number concentrations in general, thus on a longer time interval such as season
13 or year. The nucleation source had the largest effect on particle concentrations around noon and early afternoon as expected.
14 During this time interval, it became the major source of particles in the near-city background. Nucleation increased the daily
15 mean particle number concentrations on nucleation days by mean factors of 2.3 and 1.58 in the near-city background and city
16 centre, respectively. Its effect was the largest in winter, which was explained with the substantially lower background
17 concentration levels on nucleation days than that on non-nucleation days. On an annual time scale, 37% of the UF particles
18 were generated by nucleation in the near-city background, while NPF produced 13% of UF particles in the city centre. The
19 differences among the annual mean values, and among the corresponding seasonal mean values were likely caused by the
20 variability in controlling factors from year to year. The values obtained represent lower limits of contributions. The shares
21 determined imply that NPF is a non-negligible or substantial source of particles in near-city background environments and
22 even in city centres, where the vehicular road emissions usually prevail. Atmospheric residence time of nucleation-mode
23 particles was assessed by decay curve analysis of N_{6-25} concentrations in time, and a mean of 2:30 was obtained. The present
24 study suggests that the health-related consequences of atmospheric NPF and growth process in cities should also be considered
25 in addition to its urban climate implications.

26 1 Introduction

27 Large-scale modelling studies suggest that new aerosol particle formation (NPF) and consecutive particle diameter growth
28 process in the atmosphere (Kulmala et al., 2004, 2013) is the dominant source of particle number concentrations on global
29 scale (Spracklen et al., 2006; Reddington et al., 2011; Makkonen et al., 2012; Yu et al., 2015). In addition, up to approximately
30 50% of all cloud condensation nuclei (CCN) can originate from NPF and growth (Spracklen et al., 2008; Merikanto et al.,
31 2009), which relates the process to the climate system, and indicates its overall importance (Kerminen et al., 2012; Carslaw et
32 al., 2013; Shen et al., 2017). New particle formation has also been proved to be common in large cities (Nieminen et al., 2017).
33 Urban NPF can interact with and can be influenced by regional nucleation events at least under some geographic conditions,
34 and can become part of a phenomenon with a much larger horizontal extension than the city (Salma et al., 2016b). At the same
35 time, particle number concentrations in cities are strongly affected by high-temperature emission sources from different sectors
36 (Paasonen et al. 2016) such as household and residential heating (e.g. Butt et al., 2016), industrial processes and power
37 production (e.g. Xiao et al., 2015), and vehicular road traffic (e.g. Morawska et al., 2008). Their diurnal variability often show
38 daily time-activity pattern of inhabitants (Dall'Osto et al., 2013). Relative contributions of primary and secondary particle
39 sources – particularly in cities – change substantially in time and space (Pikridas et al., 2015; Posner and Pandis, 2015). Several



40 methods were proposed to distinguish the major production types of particles (e.g. Shi et al., 1999; Alam et al., 2003; Rodrigues
41 and Cuevas, 2007; Qian et al., 2007; Park et al., 2008; Costabile et al., 2009; Brines et al., 2015). The share of NPF as a single
42 source of ambient particle number concentrations specifically in cities remained, however, largely unknown. Despite the fact
43 that there is often a spatial coincidence between the poorer air quality and population density (Samoli et al., 2016). Moreover,
44 approximately 70–80% of total particles in cities belong to the ultrafine (UF) size range (with an equivalent diameter <100
45 nm; Putaud et al., 2010), and their inhalation can represent an excess health risk relative to coarse or fine particles with the
46 same or similar chemical composition (Oberdörster et al., 2005; Braakhuis et al., 2014). An estimate on the relative contribution
47 of primary and secondary formation processes is also required for efficient action plans to improve the air quality in cities. It
48 is worth noting that indirect climate effects (due to CCN) become important for particles with a diameter >50–100 nm, while
49 the excess health effects are linked with diameters <100 nm.

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51 Nucleation strength factor (NSF) was introduced to assess the contribution of NPF to UF particle number concentrations
52 relative to the background with respect to all other sources (Salma et al., 2014). The results derived from this approach
53 correspond to the mode-segregated secondary particle load. By now, atmospheric concentration data sets are available for
54 multiple years to study the applicability and behaviour of NSF in detail. The major advantage of this quantification is that it
55 only requires experimental data that can be readily derived from ordinary NPF (size distributions) measurements. The main
56 objectives of this paper are to quantify and discuss the contribution of NPF events to ambient particle number concentrations
57 in near-city and central urban environments of a Central European city considering five-year long data sets, to investigate and
58 explain the details of the NSF, and to interpret the consequences achieved for the urban air and air quality.

59 2 Methods

60 2.1 Experimental

61 The measurements were performed in Budapest, Hungary. Its population is approximately 2.5 million in the metropolitan area.
62 The major pollution sources in terms of particle number include vehicular road traffic, residential heating and household
63 burning activities. Contributions of passenger cars and buses to the vehicle fleet registered in Budapest and Pest County are
64 87% and 0.46%, respectively (OKJ, 2015). Diesel-powered vehicles shared 19% and 97% of the national passenger car and
65 bus fleets, respectively. Wintertime median concentrations of particulate matter (PM) mass, elemental carbon (EC) and organic
66 carbon (OC) in the PM_{2.5} size fraction were 25, 0.97 and 4.9 µg m⁻³, respectively in the related time interval (Salma et al.,
67 2017). The mean contributions of EC and organic matter (OM, with an OM/OC mass conversion factor of 1.6) to the PM_{2.5}
68 mass and standard deviations (SDs) were 4.8±2.1% and 37±10%, respectively, while the contribution of (NH₄)₂SO₄ and
69 NH₄NO₃ derived from an earlier study in spring were 24% and 3%, respectively. The contributions of EC and OC from fossil
70 fuel combustion to the TC were 11.0% and 25%, respectively, and EC and OC from biomass burning were responsible for
71 5.8% and 34%, respectively of the TC, while the OC from biogenic sources made up 24% of the TC.

72

73 Two urban sites were involved in the study. Most measurements were performed at the Budapest platform for Aerosol Research
74 and Training (BpART) facility (N 47° 28' 29.9", E 19° 3' 44.6", 115 m above mean sea level (a.s.l.) of the Eötvös University
75 (Salma et al., 2016a). The sampling inlets were set up at heights between 12 and 13 m above the street level. The location
76 represents a well-mixed, average atmospheric environment for the city centre. The other location was situated at the NW border
77 of Budapest in a wooded area of the Konkoly Astronomical Observatory of the Hungarian Academy of Sciences (N 47° 30'
78 00.0", E 18° 57' 46.8", 478 m a.s.l.). It represents the air masses entering the city since the prevailing wind direction in the
79 Budapest area is NW. The experimental data obtained for five full-year long time intervals, i.e. from 03–11–2008 to 02–11–
80 2009, from 19–01–2012 to 18–01–2013, from 13–11–2013 to 12–11–2014, from 13–11–2014 to 12–11–2015 and from 13–



81 11–2015 to 12–11–2016 were considered in the present study. Local time (UTC+1 and daylight saving time, UTC+2) was
 82 chosen as the time scale because the daily routine activities of inhabitants in cities were primarily considered.

83

84 The key measuring instrument was a flow-switching type differential mobility particle sizer (DMPS; Salma et al., 2011). Its
 85 main components include a Ni-60 radioactive bipolar charger, a Nafion semi-permeable membrane dryer, a 28-cm long
 86 Vienna-type differential mobility analyser and a butanol-based condensation particle counter (TSI, model 3775). The system
 87 operates in an electrical mobility diameter range from 6 to 1000 nm in the dry state of particles (with a RH<30%) in 30 channels
 88 with a time resolution of approximately 8 or 10 min at two sets of flows. The sample flow rate is 2.0 L min⁻¹ in high-flow
 89 mode, and 0.31 L min⁻¹ in low-flow mode with sheath air flow rates 10 times larger than for the sample flows. The DMPS
 90 measurements were performed according to the recommendations of the international technical standard (Wiedensohler et al.,
 91 2012). The DMPS data for the 1-year long time intervals in 2008–2009, 2012–2013, 2013–2014, 2014–2015 and 2015–2016
 92 were available in 95%, 95%, 99%, 95% and 73% of the total number of days, respectively. Meteorological data were recorded
 93 by an on-site meteorological station (Salma et al., 2016a). Standardised meteorological measurements of air temperature (*T*),
 94 relative humidity (RH), wind speed (WS) and wind direction (WD) were recorded with a time resolution of 10 min. The
 95 coverage of the meteorological data was >80% in each year.

96 2.2 Data treatment

97 The overall treatment of the measured DMPS data was performed according to the procedure protocol by Kulmala et al. (2012).
 98 The inverted DMPS data were utilised to generate particle number size distribution surface plots showing jointly the variation
 99 in particle diameter and particle number concentration density in time. Identification and classification of NPF and growth
 100 processes was accomplished from the surface plots by using the algorithm similar to that of Dal Maso et al. (2005) on a day-
 101 to-day basis into the following main classes: NPF event days, non-event days, days with undefined character, and days with
 102 missing data (for more than 4 h in the midday). Frequency of events was determined as the ratio of the number of event days
 103 to the total number of relevant (i.e. all–missing) days. Particle number concentrations in the diameter ranges from 6 to 1000
 104 nm (N_{6-1000}), from 6 to 100 nm (N_{6-100}), from 6 to 25 nm (N_{6-25}) and from 100 to 1000 nm ($N_{100-1000}$) were calculated from the
 105 DMPS data. The major portion of the N_{6-100} concentration (i.e. the Aitken mode) is essentially related to local source processes
 106 due to their limited atmospheric residence time (typically <10¹ h), while the $N_{100-1000}$ (concentration mainly in the accumulation
 107 mode) expresses larger (background) spatial and time scales because of much longer residence times (up to 10¹ d; Salma et
 108 al., 2011). To derive mean diurnal variability, the exact recording times belonging to individual concentrations were rounded
 109 off to 5 min (in case of the time resolution of ca. 8 min) or 10 min (in case of the time resolution of ca. 10 min) time scale. The
 110 concentrations and further properties derived from them (see later) were averaged by the time of day separately for nucleation
 111 and non-nucleation days, and in a year. Finally, the averaging was also performed separately for different seasons, hence for
 112 spring (March–May), summer (June–August), autumn (September–November) and winter (December–February).

113

114 Two types of NSF (Salma et al., 2014) were derived in the present study by considering different conditions. The quantity:

$$115 \text{ NSF}_{\text{nucl days}} = \frac{\left(\frac{N_{6-100}}{N_{100-1000}} \right)_{\text{nucl days}}}{\left(\frac{N_{6-100}}{N_{100-1000}} \right)_{\text{non-nucl days}}} \quad (1)$$

116 considers the $N_{6-100}/N_{100-1000}$ concentration ratios for nucleation days only. The numerator expresses the increase in N_{6-100}
 117 concentration relative to the background concentration $N_{100-1000}$ caused by all source sectors. The denominator represents the
 118 same property due to all sources except for NPF. Hence, the $\text{NSF}_{\text{nucl days}}$ accounts for the increment in background particle
 119 concentration on nucleation days exclusively caused by NPF. It was implicitly assumed that the major emission and formation



120 processes of UF particles except for NPF are uniformly present on both nucleation and non-nucleation days. It seems to be a
 121 reasonable condition for time intervals of several months, although the number of nucleation days during a time interval
 122 actually plays a more determining role than the length of the interval. Winter, when the occurrence frequency shows the
 123 minimum (see Table 1), appears to be the most restrictive season. The effect of the non-uniformly present sources is indicated
 124 by unusually larger scatter in the diurnal data points (see Sect. 3.2). It was also presumed that the production of particles larger
 125 than 100 nm was much smaller than the concentration of UF particles. This is ordinarily realised in cities, and can be justified
 126 from the contributions of UF particles to the total particle number (Putaud et al., 2010; Németh et al., 2017).

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128 The other type of NSF was calculated for all days in the numerator, thus:

$$129 \quad NSF_{\text{all days}} = \frac{\left(\frac{N_{6-100}}{N_{100-1000}} \right)_{\text{all days}}}{\left(\frac{N_{6-100}}{N_{100-1000}} \right)_{\text{non-nucleation days}}} \quad (2)$$

130 It expresses the overall contribution of NPF to particle numbers of background concentration on longer time span or in general.
 131 Since there are usually more non-nucleation days than nucleation days in a time interval of month or more, the assumptions
 132 for $NSF_{\text{all days}}$ are met easier than for $NSF_{\text{nucl days}}$. The $NSF_{\text{nucl days}}$ characterises an ordinary nucleation day within e.g. a season,
 133 while the $NSF_{\text{all days}}$ quantifies the overall effect of NPF and growth on the atmospheric concentrations on a typical or average
 134 day over e.g. a season or year. If 1) $NSF \approx 1$ then the relative contribution of nucleation to particle number concentrations with
 135 respect to other sources is negligible, 2) $1 < NSF < 2$ then its relative contribution as a single source is considerable, and 3)
 136 $NSF > 2$ then the contribution of nucleation itself to particle number concentrations is larger than of any other source sectors
 137 together. Since the major phase of NPF and growth process takes place in most cases in one day, it is advantageous to express
 138 NSF as daily mean values. The data for the undefined days were not taken into account for the present evaluation.

139 3 Results and discussion

140 Number of nucleation days for different seasons in each measurement year are summarised in Table 1. It is seen that the NPF
 141 frequency has an obvious seasonal variability. This can be obtained from its monthly dependency which exhibits an absolute
 142 and local minimum in January and August, respectively, and an absolute and local maximum in March or April, and September,
 143 respectively (Salma et al., 2016b). The seasonal variation of the nucleation frequency fits into the second group of the
 144 measurement sites reported by Manninen et al. (2010).

145 **Table 1.** Number of nucleation days for seasons in the near-city background (in 2012–2013) and in the city centre (in 2008–2009, 2013–
 146 2014, 2014–2015 and 2015–2016) during 1-year long time intervals.

Environment	Time interval	Spring	Summer	Autumn	Winter
Background	2012–2013	35	20	24	17
Centre	2008–2009	34	21	22	6
Centre	2013–2014	28	20	13	11
Centre	2014–2015	41	19	14	7
Centre	2015–2016	15	9*	5*	6

147 * Low data coverage. See Sect. 2.1.



148 3.1 Seasonal atmospheric concentrations

149 Particle number concentrations in the related size fractions for different seasons in each year are summarised in Table 2 for an
 150 overview.

151 **Table 2.** Median atmospheric concentration of particles with a diameter from 6 to 100 nm (N_{6-100}) and from 100 to 1000 nm ($N_{100-1000}$) in
 152 units of 10^3 cm^{-3} separately on nucleation (Nucl) days and non-nucleation (Nonucl) days for seasons in the near-city background (in 2012–
 153 2013) and in the city centre (in 2008–2009, 2013–2014, 2014–2015 and 2015–2016) during 1-year long time intervals. Mean ratios of median
 154 concentrations on nucleation days to that on non-nucleation days with standard deviations (SDs) for the size fraction are also indicated.

Urban environment			Near-city backgr.		City centre				
Season	Size fraction	Day type	2012–2013	Ratio	2008–2009	2013–2014	2014–2015	2015–2016	Ratio \pm SD
Spring	N_{6-100}	Nucl	4.8	1.72	11.2	9.7	10.0	8.6	1.37
	N_{6-100}	Nonucl	2.8		8.9	7.2	7.1	5.9	± 0.09
	$N_{100-1000}$	Nucl	1.56	1.03	2.0	2.5	2.6	1.56	0.99
	$N_{100-1000}$	Nonucl	1.51		2.2	2.7	2.5	1.5	± 0.07
Summer	N_{6-100}	Nucl	4.0	1.37	10.3	8.0	8.6	6.9	1.17
	N_{6-100}	Nonucl	2.9		8.9	7.5	6.5	6.2	± 0.11
	$N_{100-1000}$	Nucl	1.27	0.89	1.36	2.0	2.5	1.40	0.92
	$N_{100-1000}$	Nonucl	1.42		1.72	2.4	2.4	1.37	± 0.13
Autumn	N_{6-100}	Nucl	4.3	1.29	14.0	11.9	12.6	5.2	1.41
	N_{6-100}	Nonucl	3.3		10.4	8.5	8.4	5.1	± 0.08
	$N_{100-1000}$	Nucl	1.67	0.74	2.0	3.4	2.7	1.6	0.87
	$N_{100-1000}$	Nonucl	2.3		2.4	3.9	3.3	1.7	± 0.06
Winter	N_{6-100}	Nucl	3.9	1.10	6.9	10.5	5.6	7.7	0.87
	N_{6-100}	Nonucl	3.6		12.5	9.2	7.8	7.4	± 0.28
	$N_{100-1000}$	Nucl	1.12	0.38	1.02	3.7	1.65	1.4	0.54
	$N_{100-1000}$	Nonucl	2.9		3.0	4.5	3.8	2.7	± 0.22

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156 Data coverage for summer and autumn in 2015–2016 were low, and therefore, the corresponding concentration ratios were
 157 excluded from the averaging for the mean ratios. It can be seen that the N_{6-100} were ordinarily larger on nucleation days than
 158 on non-nucleation days. This is likely a direct effect of nucleation. At the same time, the $N_{100-1000}$ usually showed a constant
 159 level within approximately 10% except for winters and some autumns. The background concentrations on nucleation days
 160 were smaller (by factors of 0.4–0.5) than for non-nucleation days particularly in winter. The differences are further discussed
 161 and explained in Sect. 3.2.

162 3.2 Concentration increment on nucleation days

163 Diurnal variability of the concentration increment due to NPF on nucleation days (i.e. $\text{NSF}_{\text{nucl days}}$) for the city centre and near-
 164 city background separately for different seasons are shown in Fig. 1 as representative examples. The curves exhibited a single
 165 peak around noon with a longer tail on the decreasing side. The exact location of the peak is also influenced by setting the
 166 local daylight saving time in spring and autumn. The baseline of some peaks from 0:00 to 7:00 deviated systematically and
 167 substantially from unity although no nocturnal nucleation has been observed in Budapest. The mean values of this baseline in
 168 the city centre for spring, summer, autumn and winter (Fig. 1 upper panel) were 1.02, 1.15, 1.15 and 1.55, respectively, while
 169 they were 1.11, 1.18, 1.31 and 1.72, respectively in the near-city background (Fig. 1 lower panel). This elevated line can be
 170 explained by the fact that particle growth process could be traced till the late morning of the next day in several occasions,
 171 thus the NPF influenced the N_{6-100} concentrations over the next morning. This affected the baseline if a non-nucleation day
 172 followed a nucleation day, and particularly, in the seasons when NPF events occur well separated from each other in time,
 173 which is typical for winter. The elevated line is a real effect of NPF, and its consideration in the averaging is justified.

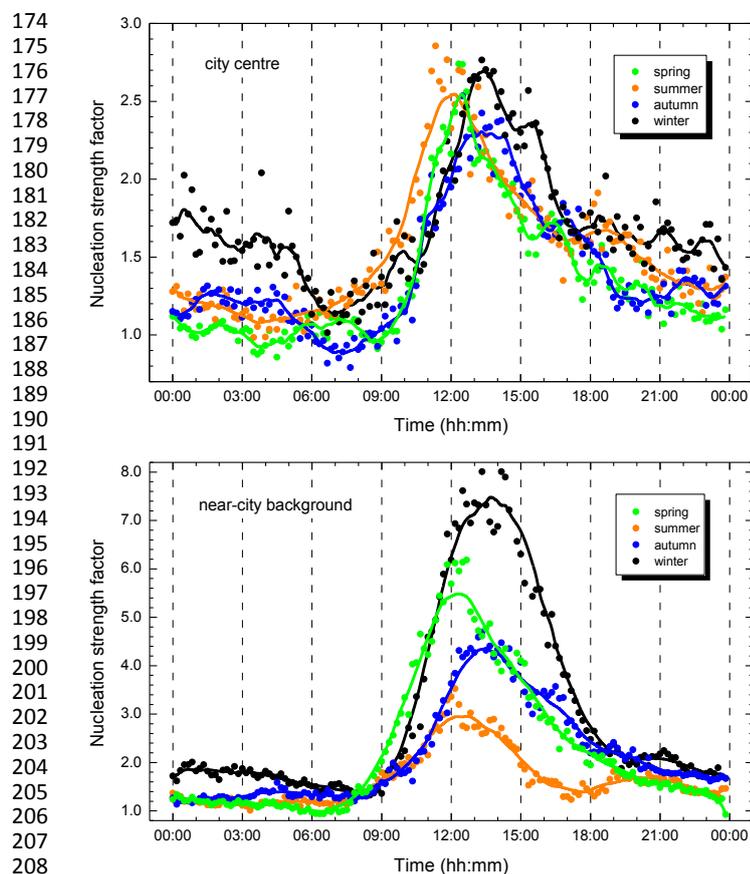


Figure 1. Diurnal variability of concentration increment due to NPF on nucleation days (nucleation strength factor $NSF_{nucl\ days}$) for the city centre in 2008–2009 (upper panel) and in the near-city background in 2012–2013 (lower panel) separately for seasons. The solid lines represent 1-h smoothing to the data.

It was also observed in all years that the concentration increment on nucleation days due to NPF (i.e. $NSF_{nucl\ days}$) was the largest for winter. This evidently showed up for the near-city background. It was followed by the other seasons which had similar importance to each other in the city centre, or which were ordered as spring, autumn and summer in the near-city background. To investigate these findings more closely, diurnal variability of the related particle number concentrations were derived and evaluated. Diurnal variability of the concentrations for summer and winter are shown in Figs. 2 and 3, respectively for the city centre and near-city background. The dependencies in the city centre for the spring, other summer and autumn seasons are similar to Fig. 2 upper panel, while the corresponding seasonal curves for the near-city background resemble Fig. 2 lower panel. The diurnal patterns represented by Fig. 2 are coherent with the previous ideas on the NPF and growth events in the Budapest area (Salma et al., 2014, 2016b; Németh et al., 2017), and they also confirm the basic assumptions of the NSF definition on the source intensities at both location types. The comparison of the N_{6-100} curves for nucleation days and non-nucleation days already emphasizes the importance of NPF, and indicates that the phenomenon has larger relative effect in the near-city background than in the central urban parts as it is expected. It is worth realising that the particle number concentrations for the background aerosol ($N_{100-1000}$) appear close to each other within a relative uncertainty of 10–20%, which implies that the background concentrations affect the NPF occurrence and formation rate in a limited manner in these seasons.

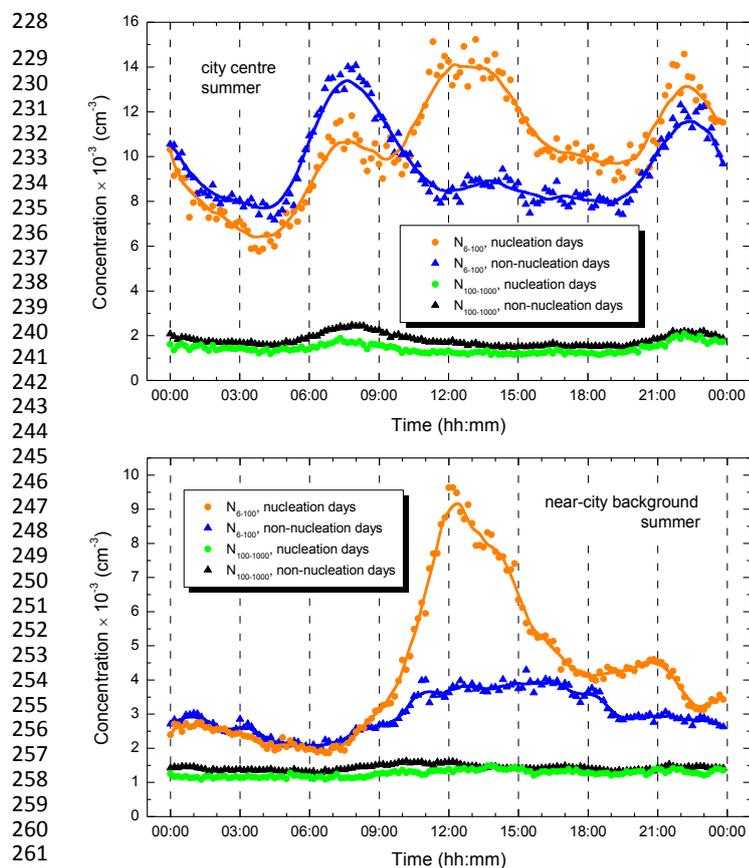


Figure 2. Diurnal variability of particle number concentrations in summer for the diameter ranges from 6 to 100 nm (N_{6-100}) and from 100 to 1000 nm ($N_{100-1000}$) in the city centre in 2008–2009 (upper panel) and in the near-city background in 2012–2013 (lower panel) separately for nucleation days and non-nucleation days. The solid lines represent 1-h smoothing to the data.

267 For winter, it is seen, however, that the background concentrations were substantially different for the nucleation and non-
 268 nucleation days (Fig. 3). The mean non-nucleation/nucleation $N_{100-1000}$ ratios for the city centre (in 2008–2009) and near-city
 269 background were 2.8 and 2.3, respectively. This implies that NPF events preferably took place on those days when the particle
 270 number concentrations were generally smaller. It is understandable if we consider that the basic preconditions of NPF events
 271 are realised by competing source and sink for condensing vapours. The source strength in winter has a decreased tendency due
 272 to lower solar radiation intensities and less (biogenic) precursor gases in the air (Salma et al., 2011). Nevertheless, nucleation
 273 can occur at these small source terms if the (condensation and scavenging) sink - which is related to the concentration of pre-
 274 existing particles - is even smaller. This explains the differences in the background concentrations on nucleation and non-
 275 nucleation days. Larger concentration increments (higher $NSF_{nucl\ days}$, Fig. 3) for winter were simply caused by systematically
 276 smaller background concentrations on nucleation days. In addition, the fluctuation in the $N_{100-1000}$ for nucleation days in winter
 277 of the other years was sometimes larger than that shown in Fig. 3. This observation raises the question which is the smallest
 278 number of NPF events in a time interval (e.g. season), which can be considered to be sufficient for obtaining representative
 279 mean diurnal concentration data for calculating the NSF. A few NPF events during winters (see Table 1) might not be fully
 280 satisfactory for this purpose.

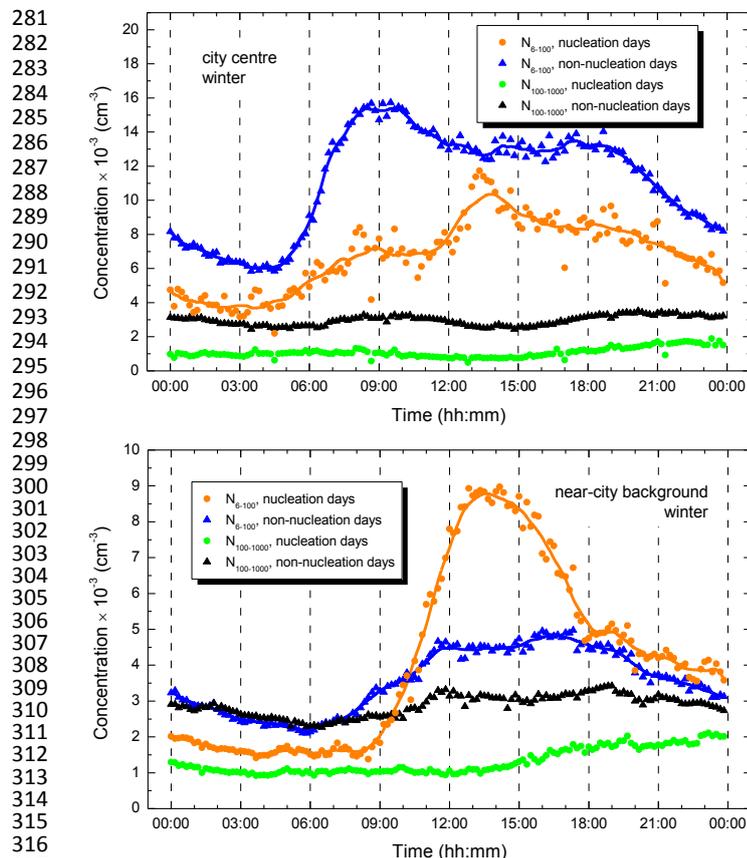


Figure 3. Diurnal variability of particle number concentrations in winter for the diameter ranges from 6 to 100 nm (N_{6-100}) and from 100 to 1000 nm ($N_{100-1000}$) in the city centre in 2008–2009 (upper panel) and in the near-city background in 2012–2013 (lower panel) separately for nucleation days and non-nucleation days. The solid lines represent 1-h smoothing to the data.

Seasonal and annual mean concentration increments of background concentration on nucleation days (hence the daily mean $NSF_{nuc\ days}$ values) for different years are summarised in Table 3. Nucleation as a single source increased the daily particle number concentrations by factors of 2.3 and 1.58 in the near-city background and city centre, respectively on an annual time scale. The differences among the annual mean values, and among the corresponding seasonal mean values were likely caused by the variability of controlling parameters from a year to year. As far as the seasonal variability is concerned, it is noted that the formation rate for particles with a diameter of 6 nm (J_6) showed only rather modest seasonal dependency in Budapest (Salma et al., 2011), which also contributed to similar mean increments for spring, summer and autumn.

Table 3. Seasonal and annual mean increments of background concentration due to nucleation on nucleation days (nucleation strength factor $NSF_{nuc\ days}$) in the near-city background (in 2012–2013) and in the city centre (in 2008–2009, 2013–2014, 2014–2015 and 2015–2016) during 1-year long time intervals.

Environment	Time interval	Spring	Summer	Autumn	Winter	Year
Background	2012–2013	2.3	1.66	2.2	3.0	2.3
Centre	2008–2009	1.36	1.55	1.42	1.71	1.49
Centre	2013–2014	1.31	1.33	1.36	1.56	1.44
Centre	2014–2015	1.50	1.34	1.83	2.8	1.73
Centre	2015–2016	1.54	1.26	1.46	2.4	1.64

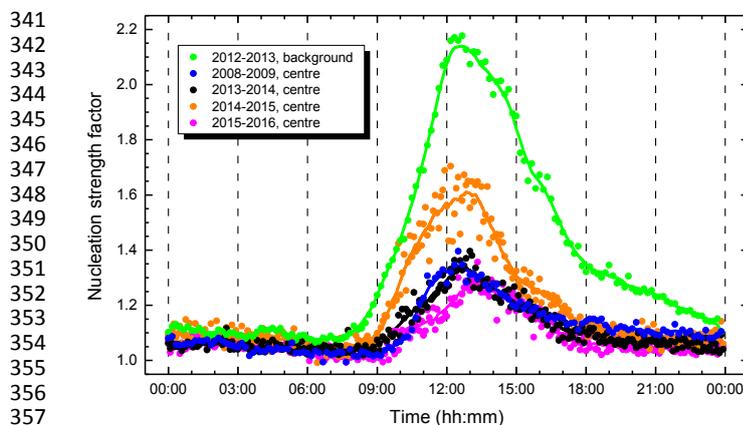
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333 3.3 Contribution of nucleation to particle number concentrations

334 Diurnal variability of $NSF_{all\ days}$ is shown in Fig. 4. The curves exhibited a single peak with a maximum around noon and a
335 longer tail in the early afternoon. The maximum values in the city centre represented concentration contributions from 30% to
336 60% due to nucleation for a limited time interval. The curve for the near-city background was the largest, as expected, and it
337 even exceeded the value of 2 around noon for approximately 3 h. This all means that nucleation has an important contribution
338 to UF particles during the midday in the city centre, while it even becomes the dominant source of particles directly after
339 midday in the near-city background.

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359 **Figure 4.** Diurnal variability of NPF contribution to particle number concentrations (nucleation strength factor $NSF_{all\ days}$) in the near-city
360 background (in 2012–2013) and in the city centre (in 2008–2009, 2013–2014, 2014–2015 and 2015–2016) during 1-year long time intervals.
361 The solid lines represent 1-h smoothing to the data.
362

363 The importance of nucleation was also demonstrated for different seasons and years by the mean $NSF_{all\ days}$ values which are
364 summarised in Table 4. In general, 37% of UF particles (more precisely of particle number increment) were produced by
365 nucleation as a single source in the near-city background. In the city centre, it generated 13% of UF particles. These values
366 can be considered as lower limits since a considerable part of the background particles is also produced by NPF. The differences
367 among the annual mean values, and among the corresponding seasonal mean values were likely caused by the year-to-year
368 variability similarly to the concentration increments on nucleation days ($NSF_{nucl\ days}$). It is informative to compare the
369 contribution values to the global share of various source sectors in primary UF particle number emission to have an idea on
370 the relative extent of our results. It is stressed that our and the literature data are related to very different types of particles,
371 nevertheless, some analogy can be found in the relative importance of the sources and source processes. Road transport, power
372 production and residential combustion are the first three largest contributors to primary UF particles with the shares of 40%,
373 20% and 17%, respectively (Paasonen et al., 2016). The actual contributions can vary in different parts of the world and with
374 economic development.

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376 The effects of particles generated by NPF and growth in increased concentrations on human health and the environment are
377 also influenced by the time interval for which the particles remain in the air. Nucleation-mode particles are primarily removed
378 by coagulation with larger particles, agglomeration, diffusion and turbulent losses, scavenging and various aging processes.
379 We showed in Sect. 3.2 and it can be also proved directly from the measured data that it is the NPF events that usually produce
380 the N_{6-25} concentrations in an overwhelming extent (the N_{6-25} is increased by 1–2 orders of magnitude) in a relatively short
381 time interval. The continual concentration decrease in time for several hours after the event can be utilised to assess the general
382 atmospheric residence time of nucleation-mode particles. The decrease of this concentration after a nucleation burst could be



383 approximated by an exponential function (first order kinetics). By a decay curve analysis of 15 selected NPF and particle
384 diameter growth cases, the residence times were estimated from the slope of the concentrations N_{6-25} in a natural logarithmic
385 scale versus time. The residence time of nucleated particles varied from 1:30 to 4:15 with a mean and SD of $2:30 \pm 1:00$. This
386 suggests that the nucleation-mode particles (or in other words the nucleated particles in their very small sizes, or atmospheric
387 nanoparticles) likely have limited health effects due to their relatively short existence in the air.

388 **Table 4.** Seasonal and annual mean contributions of nucleation to background concentration (nucleation strength factor $NSF_{all\ days}$) in the
389 near-city background (in 2012–2013) and in the city centre (in 2008–2009, 2013–2014, 2014–2015 and 2015–2016) during 1-year long time
390 intervals.

Environment	Time interval	Spring	Summer	Autumn	Winter	Year
Background	2012–2013	1.51	1.18	1.31	1.40	1.37
Centre	2008–2009	1.12	1.15	1.10	1.05	1.12
Centre	2013–2014	1.11	1.07	1.08	1.11	1.11
Centre	2014–2015	1.22	1.08	1.14	1.16	1.19
Centre	2015–2016	1.09	1.06	1.03	1.09	1.09

391 4 Conclusions

392 We showed in the present study that NPF and particle diameter growth process as a single source represents a considerable
393 contribution to UF particles in a Central European city with respect to all other emission sources including vehicular road
394 traffic. Nucleation was a major process that produced UF particles at noon and in the early afternoon, and its relative
395 contribution was comparable to other production sources during this time period even in the city centre. Relative importance
396 of nucleation as a source of particles decreased with anthropogenic influence. The NSF_s were defined by utilising N_{6-100} and
397 $N_{100-1000}$ concentrations. There are several sensible and practical reasons for selecting these specific size fractions although
398 other dividing values are also imaginable. The quantifications in the present study are, therefore, subjected to certain inherent
399 uncertainty. Considering that the modes of the particle number size distribution are usually shifted to smaller diameters in
400 cities with respect to rural or remote areas, it seems realistic that these size fractions represent well the particles of urban (local)
401 origin and the aged particles (which characterize larger spatial or urban background area), respectively. The study also suggests
402 that particles from NPF events in cities are relevant not only for their effects on urban climate but because of their health risk
403 to inhabitants. At the same time, it should also be mentioned that ambient atmospheric aerosol which ordinary persists in the
404 air of cities contains particles in the largest abundance with a diameter between approximately 25 and 150 nm. Smaller particles
405 are thermodynamically not stable, and most of them are removed from the air during relatively short time intervals. The
406 exposures to freshly nucleated particles ($d < 25$ nm) or ambient nanoparticles ($d < 10$ nm) are usually limited to several hours
407 after the onset of the NPF.

408
409 Regulations of aerosol emissions and atmospheric concentrations are usually based on PM mass. The changes or reductions in
410 anthropogenic aerosol load are ordinary assessed by assuming similar relative tendencies in particle mass and particle number
411 concentrations. This generalisation may yield to tentative conclusions. According to current legislation scenarios, particle
412 number emissions are expected to decrease in most part of the world by 2030 mainly due to spreading the diesel particulate
413 filters (DPF) in cars and due to diesel fuels with ultralow sulphur content. In effect, this may imply that the relative share of
414 NPF in the particle number production is expected to increase above the levels estimated in the present study. By demonstrating
415 the relevance of NPF as an important single source of UF particles, we also raise the question of an international enhanced
416 particle mass and particle number inventory with precursor gas data that potentially includes the NPF and growth process as a
417 separate sector among the source types.

418 **5 Data availability**

419 The relevant observational data used in this paper are available on request from the corresponding author or at the website of
420 the Budapest platform for Aerosol Research and Training (<http://salma.web.elte.hu/BpArt>).

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423 **References**

- 424 Alam, A., Shi, J. P., and Harrison, R. M.: Observation of new particle formation in urban air, *J. Geophys. Res.*, 108, 4093–
425 4107, 2003.
- 426 Braakhuis, H. M., Park, M. V., Gosens, I., De Jong, W. H., and Cassee, F. R.: Physicochemical characteristics of
427 nanomaterials that affect pulmonary inflammation, *Part. Fibre Toxicol.*, 11:18, doi: 10.1186/1743-8977-11-18, 2014.
- 428 Brines, M., Dall'Osto, M., Beddows, D. C. S., Harrison, R. M., Gómez-Moreno, F., Núñez, L., Artíñano, B., Costabile, F.,
429 Gobbi, G. P., Salimi, F., Morawska, L., Sioutas, C., and Querol, X.: Traffic and nucleation events as main sources of
430 ultrafine particles in high-insolation developed world cities, *Atmos. Chem. Phys.*, 15, 5929–5945, 2015.
- 431 Butt, E. W., Rap, A., Schmidt, A., Scott, C. E., Pringle, K. J., Reddington, C. L., Richards, N. A. D., Woodhouse, M. T.,
432 Ramirez-Villegas, J., Yang, H., Vakkari, V., Stone, E. A., Rupakheti, M., S. Praveen, P., G. van Zyl, P., P. Beukes, J.,
433 Josipovic, M., Mitchell, E. J. S., Sallu, S. M., Forster, P. M., and Spracklen, D. V.: The impact of residential combustion
434 emissions on atmospheric aerosol, human health, and climate, *Atmos. Chem. Phys.*, 16, 873–905, 2016.
- 435 Carslaw, K. S., Lee, L. A., Reddington, C. L., Pringle, K. J., Rap, A., Forster, P. M., Mann, G. W., Spracklen, D. V.,
436 Woodhouse, M. T., Regayre, L. A., and Pierce, J. R.: Large contribution of natural aerosols to uncertainty in indirect
437 forcing, *Nature*, 503, 67–71, 2013.
- 438 Carslaw, K. S., Lee, L. A., Reddington, C. L., Pringle, K. J., Rap, A., Forster, P. M., Mann, G. W., Spracklen, D. V.,
439 Woodhouse, M. T., Regayre, L. A., and Pierce, J. R.: Large contribution of natural aerosols to uncertainty in indirect
440 forcing, *Nature*, 503, 67–71, 2013.
- 441 Costabile, F., Birmili, W., Klose, S., Tuch, T., Wehner, B., Wiedensohler, A., Franck, U., König, K., and Sonntag, A.:
442 Spatio-temporal variability and principal components of the particle number size distribution in an urban atmosphere,
443 *Atmos. Chem. Phys.*, 9, 3163–3195, 2009.
- 444 Dal Maso, M., Kulmala, M., Riipinen, I., Wagner, R., Hussein, T., Aalto, P. P., and Lehtinen, K. E. J.: Formation and growth
445 of fresh atmospheric aerosols: eight years of aerosol size distribution data from SMEAR II, Hyytiälä, Finland, *Boreal*
446 *Environ. Res.*, 10, 323–336, 2005.
- 447 Dall'Osto, M., Querol, X., Alastuey, A., O'Dowd, C., Harrison, R. M., Wenger, J., and Gómez-Moreno, F. J.: On the spatial
448 distribution and evolution of ultrafine particles in Barcelona, *Atmos. Chem. Phys.*, 13, 741–759, 2013.
- 449 Kerminen, V.-M., Paramonov, M., Anttila, T., Riipinen, I., Fountoukis, C., Korhonen, H., Asmi, E., Laakso, L., Lihavainen,
450 H., Swietlicki, E., Svenningsson, B., Asmi, A., Pandis, S. N., Kulmala, M., and Petäjä, T.: Cloud condensation nuclei
451 production associated with atmospheric nucleation: a synthesis based on existing literature and new results, *Atmos.*
452 *Chem. Phys.*, 12, 12037–12059, 2012.
- 453 Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V.-M., Birmili, W., and McMurry, P.:
454 Formation and growth rates of ultrafine atmospheric particles: a review of observations, *J. Aerosol Sci.*, 35, 143–176,
455 2004.
- 456 Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M., Aalto, P. P., Junninen, H.,
457 Paasonen, P., Riipinen, I., Lehtinen, K. E. J., Laaksonen, A., and Kerminen, V.-M.: Measurement of the nucleation of
458 atmospheric aerosol particles, *Nature Protocols*, 7, 1651–1667, doi:10.1038/nprot.2012.091, 2012.
- 459 Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petäjä, T., Sipilä, M.,
460 Schobesberger, S., Rantala, P., Franchin, A., Jokinen, T., Järvinen, E., Äijälä, M., Kangasluoma, J., Hakala, J., Aalto,



- 461 P.P., Paasonen, P., Mikkilä, J., Vanhanen, J., Aalto, J., Hakola, H., Makkonen, U., Ruuskanen, T., Mauldin, R. L. III,
462 Duplissy, J., Vehkamäki, H., Bäck, J., Kortelainen, A., Riipinen, I., Kurtén, T., Johnston, M. V., Smith, J. N., Ehn, M.,
463 Mentel, T. F., Lehtinen, K. E. J., Laaksonen, A., Kerminen, V.-M., and Worsnop, D. R.: Direct observations of
464 atmospheric aerosol nucleation, *Science*, 339, 943–946, 2013.
- 465 Makkonen, R., Asmi, A., Kerminen, V.-M., Boy, M., Arneth, A., Hari, P., and Kulmala, M.: Air pollution control and
466 decreasing new particle formation lead to strong climate warming, *Atmos. Chem. Phys.*, 12, 1515–1524, 2012.
- 467 Manninen, H. E., Nieminen, T., Asmi, E., Gagné, S., Häkkinen, S., Lehtipalo, K., Aalto, P., Vana, M., Mirme, A., Mirme, S.,
468 Hörrak, U., Plass-Dülmer, C., Stange, G., Kiss, G., Hoffer, A., Törö, N., Moerman, M., Henzing, B., de Leeuw, G.,
469 Brinkenberg, M., Kouvarakis, G. N., Bougiatioti, A., Mihalopoulos, N., O'Dowd, C., Ceburnis, D., Arneth, A.,
470 Svenningsson, B., Swietlicki, E., Tarozzi, L., Decesari, S., Facchini, M. C., Birmili, W., Sonntag, A., Wiedensohler, A.,
471 Boulon, J., Sellegri, K., Laj, P., Gysel, M., Bukowiecki, N., Weingartner, E., Wehrle, G., Laaksonen, A., Hamed, A.,
472 Joutsensaari, J., Petäjä, T., Kerminen, V.-M., and Kulmala, M.: EUCAARI ion spectrometer measurements at 12
473 European sites - analysis of new-particle formation events, *Atmos. Chem. Phys.*, 10, 7907–7927, 2010.
- 474 Merikanto, J., Spracklen, D. V., Mann, G. W., Pickering, S. J., and Carslaw, K. S.: Impact of nucleation on global CCN,
475 *Atmos. Chem. Phys.*, 9, 8601–8616, 2009.
- 476 Morawska, L., Ristovski, Z., Jayaratne, E. R., Keogh, D. U., and Ling, X.: Ambient nano and ultrafine particles from motor
477 vehicle emissions: characteristics, ambient processing and implications on human exposure, *Atmos. Environ.*, 42, 8113–
478 8138, 2008.
- 479 Németh, Z., Rosati, B., Ziková, N., Salma, I., Bozó, L., Dameto de España, C., Schwarz, J., Ždimal, V., and Wonaschütz, A.:
480 Comparison of atmospheric new particle formation and growth events in three Central European cities, submitted in
481 2017.
- 482 Nieminen, T., Kerminen, V.-M., Petäjä, T., Manninen, H. E., Aalto, P. P., Arshinov, M., Asmi, E., Baltensberger, U.,
483 Beukes, J. P., Collins, D., Harrison, R. M., Henzing, B., Hooda, R., Hu, M., Hörrak, U., Kivekäs, N., Komsaare, K.,
484 Krejčí, R., Laakso, L., Laaksonen, A., Leitch, R., Lihavainen, H., Mihalopoulos, N., Németh, Z., O'Dowd, C., Salma, I.,
485 Sellegri, K., Svenningsson, B., Swietlicki, E., Tunved, P., Ulevicius, V., Vakkari, V., Vana, M., Virtanen, A.,
486 Wiedensohler, A., and Kulmala, M.: Global analysis of continental boundary layer new particle formation based on long-
487 term measurements, *Atmos. Chem. Phys. Discuss.*, submitted in 2017.
- 488 Oberdörster, G., Oberdörster, E., and Oberdörster, J.: Nanotoxicology: an emerging discipline evolving from studies of
489 ultrafine particles, *Environ. Health Perspect.*, 113, 823–839, 2005.
- 490 OKJ (National register of road vehicles, in Hungarian), 2015. Ministry of National Development, Budapest.
- 491 Paasonen, P., Kupiainen, K., Klimont, Z., Visschedijk, A., Denier van der Gon, H. A. C., and Amann, M.: Continental
492 anthropogenic primary particle number emissions, *Atmos. Chem. Phys.*, 16, 6823–6840, 2016.
- 493 Park, K., Park, J. Y., Kwak, J.-H., Cho, G. N., and Kim, J.-S.: Seasonal and diurnal variations of ultrafine particle
494 concentration in urban Gwangju, Korea: Observation of ultrafine particle events, *Atmos. Environ.*, 42, 788–799, 2008.
- 495 Pikridas, M., Sciare, J., Freutel, F., Crumeyrolle, S., von der Weiden-Reinmüller, S.-L., Borbon, A., Schwarzenboeck, A.,
496 Merkel, M., Crippa, M., Kostenidou, E., Psychoudaki, M., Hildebrandt, L., Engelhart, G. J., Petäjä, T., Prévôt, A. S. H.,
497 Drewnick, F., Baltensperger, U., Wiedensohler, A., Kulmala, M., Beekmann, M., and Pandis, S. N.: In situ formation and
498 spatial variability of particle number concentration in a European megacity, *Atmos. Chem. Phys.*, 15, 10219–10237,
499 2015.
- 500 Posner, L. N. and Pandis, S. N.: Sources of ultrafine particles in the Eastern United States, *Atmos. Environ.*, 111, 103–112,
501 2015.
- 502 Putaud, J.-P., Van Dingenen, R., Alastuey, A., Bauer, H., Birmili, W., Cyrys, J., Flentje, H., Fuzzi, S., Gehrig, R., Hansson,
503 H. C., Harrison, R. M., Herrmann, H., Hitznerberger, R., Hüglin, C., Jones, A. M., Kasper-Giebl, A., Kiss, G., Kousa, A.,
504 Kuhlbusch, T. A. J., Löschau, G., Maenhaut, W., Molnár, A., Moreno, T., Pekkanen, J., Perrino, C., Pitz, M., Puxbaum,
505 H., Querol, X., Rodriguez, S., Salma, I., Schwarz, J., Smolik, J., Schneider, J., Spindler, G., ten Brink, H., Tursic, J.,
506 Viana, M., Wiedensohler, A., and Raes, F.: A European Aerosol Phenomenology - 3: physical and chemical
507 characteristics of particulate matter from 60 rural, urban, and kerbside sites across Europe, *Atmos. Environ.*, 44, 1308–
508 1320, 2010.



- 509 Reddington, C. L., Carslaw, K. S., Spracklen, D. V., Frontoso, M. G., Collins, L., Merikanto, J., Minikin, A., Hamburger, T.,
510 Coe, H., Kulmala, M., Aalto, P., Flentje, H., Plass-Dülmer, C., Birmili, W., Wiedensohler, A., Wehner, B., Tuch, T.,
511 Sonntag, A., O'Dowd, C. D., Jennings, S. G., Dupuy, R., Baltensperger, U., Weingartner, E., Hansson, H.-C., Tunved, P.,
512 Laj, P., Sellegri, K., Boulon, J., Putaud, J.-P., Gruening, C., Swietlicki, E., Roldin, P., Henzing, J. S., Moerman, M.,
513 Mihalopoulos, N., Kouvarakis, G., Ždímal, V., Zíková, N., Marinoni, A., Bonasoni, P., and Duchi, R.: Primary versus
514 secondary contributions to particle number concentrations in the European boundary layer, *Atmos. Chem. Phys.*, 11,
515 12007–12036, 2011.
- 516 Rodríguez, S. and Cuevas, E.: The contributions of “minimum primary emissions” and “new particle formation
517 enhancements” to the particle number concentration in urban air, *J. Aerosol Sci.*, 38, 1207–1219, 2007.
- 518 Salma, I., Borsós, T., Weidinger, T., Aalto, P., Hussein, T., Dal Maso, M., and Kulmala, M.: Production, growth and
519 properties of ultrafine atmospheric aerosol particles in an urban environment, *Atmos. Chem. Phys.*, 11, 1339–1353, 2011.
- 520 Salma, I., Borsós, T., Németh, Z., Weidinger, T., Aalto, P., and Kulmala, M.: Comparative study of ultrafine atmospheric
521 aerosol within a city, *Atmos. Environ.*, 92, 154–161, 2014.
- 522 Salma, I., Németh, Z., Weidinger, T., Kovács, B., and Kristóf, G.: Measurement, growth types and shrinkage of newly
523 formed aerosol particles at an urban research platform, *Atmos. Chem. Phys.*, 16, 7837–7851, 2016a.
- 524 Salma, I., Németh, Z., Kerminen, V.-M., Aalto, P., Nieminen, T., Weidinger, T., Molnár, Á., Imre, K., and Kulmala, M.:
525 Regional effect on urban atmospheric nucleation, *Atmos. Chem. Phys.*, 16, 8715–8728, 2016b.
- 526 Salma, I., Németh, Z., Weidinger, T., Maenhaut, W., Claeys, M., Molnár, M., Major, I., Ajtai, T., Utry, N., and Bozóki, Z.:
527 Source apportionment of carbonaceous chemical species to fossil fuel combustion, biomass burning and biogenic
528 emissions by a coupled radiocarbon-levoglucosan marker method, *Atmos. Chem. Phys. Discuss.*, doi:10.5194/acp-2017-
529 406, in review, 2017.
- 530 Samoli, E., Andersen, Z. J., Katsouyanni, K., Hennig, F., Kuhlbusch, T. A. J., Bellander, T., Cattani, G., Cyrus, J.,
531 Forastiere, F., Jacquemin, B., Kulmala, M., Lanki, T., Loft, S., Massling, A., Tobias, A., and Stafoggia, M.: Exposure to
532 ultrafine particles and respiratory hospitalisations in five European cities, *Eur. Resp. J.*, 48, 674–682, 2016.
- 533 Shen, L., Mickley, L. J., and Murray, L. T.: Influence of 2000–2050 climate change on particulate matter in the United
534 States: results from a new statistical model, *Atmos. Chem. Phys.*, 17, 4355–4367, 2017.
- 535 Shi, J. P., Khan, A. A., and Harrison, R. M.: Measurements of ultrafine particle concentration and size distribution in the
536 urban atmosphere, *Sci. Total Environ.*, 235, 51–64, 1999.
- 537 Spracklen, D. V., Carslaw, K. S., Kulmala, M., Kerminen, V.-M., Mann, G. W., and Sihto, S.-L.: The contribution of
538 boundary layer nucleation events to total particle concentrations on regional and global scales, *Atmos. Chem. Phys.*, 6,
539 5631–5648, 2006.
- 540 Spracklen, D. V., Carslaw, K. S., Kulmala, M., Kerminen, V.-M., Sihto, S.-L., Riipinen, I., Merikanto, J., Mann, G. W.,
541 Chipperfield, M. P., Wiedensohler, A., Birmili, W., and Lihavainen, H.: Contribution of particle formation to global
542 cloud condensation nuclei concentrations, *Geophys. Res. Lett.*, 35, L06808, doi:10.1029/2007GL033038, 2008.
- 543 Wiedensohler, A., Birmili, W., Nowak, A., Sonntag, A., Weinhold, K., Merkel, M., Wehner, B., Tuch, T., Pfeifer, S., Fiebig,
544 M., Fjåraa, A.M., Asmi, E., Sellegri, K., Depuy, R., Venzac, H., Villani, P., Laj, P., Aalto, P., Ogren, J.A., Swietlicki, E.,
545 Williams, P., Roldin, P., Quincey, P., Hüglin, C., Fierz-Schmidhauser, R., Gysel, M., Weingartner, E., Riccobono, F.,
546 Santos, S., Grünig, C., Faloon, K., Beddows, D., Harrison, R., Monahan, C., Jennings, S. G., O'Dowd, C. D., Marinoni,
547 A., Horn, H.-G., Keck, L., Jiang, J., Scheckman, J., McMurry, P. H., Deng, Z., Zhao, C. S., Moerman, M., Henzing, B.,
548 de Leeuw, G., Löschan, G., and Bastian, S.: Mobility particle size spectrometers: harmonization of technical standards
549 and data structure to facilitate high quality long-term observations of atmospheric particle number size distributions,
550 *Atmos. Meas. Tech.*, 5, 657–685, 2012.
- 551 Xiao, S., Wang, M. Y., Yao, L., Kulmala, M., Zhou, B., Yang, X., Chen, J. M., Wang, D. F., Fu, Q. Y., Worsnop, D. R., and
552 Wang, L.: Strong atmospheric new particle formation in winter in urban Shanghai, China, *Atmos. Chem. Phys.*, 15,
553 1769–1781, 2015.
- 554 Yu, F., Luo, G., Pryor, S. C., Pillai, P. R., Lee, S. H., Ortega, J., Schwab, J. J., Hallar, A. G., Leitch, W. R., Aneja, V. P.,
555 Smith, J. N., Walker, J. T., Hogrefe, O., and Demerjian, K. L.: Spring and summer contrast in new particle formation
556 over nine forest areas in North America, *Atmos. Chem. Phys.*, 15, 13993–14003, 2015.