- 1 acp-2017-817
- 2 Monoterpene chemical speciation in the Amazon tropical rainforest: variation with season, height, and
- 3 time of day at the Amazon Tall Tower Observatory (ATTO)
- 4 Combined document with detailed responses to the reviewer and a copy of the revised manuscript with5 all changes tracked.
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- 8 Response to reviewers 1 and 2.
- 9 1- Comments from reviewers
- 10 2- Author's response
- 11 3- Change in the text
- 12 Reviewer #1 (Remarks to the Author):
- 13 General:

14 The paper presents in-canopy, speciated BVOC measurements by GC-FID over a period of three days. These are compared PTR measurements and to longer time series of GC measurements in different sea-15 sons at the same site. The data is analyzed and discussed in terms of emission height, diurnal cycle, re-16 activity, and seasonality in comparison to the longer data series in different seasons. Observations are 17 compared to a canopy model which contained detailed chemistry for some species and rudimentary 18 chemistry for new others. The purpose of the comparison is to demonstrate speciated measurements are 19 20 a good test if the overall key features and processes are captured by the model. The manuscript presents interesting data and data analysis. The text is well structured. The manuscript could be published in 21 22 ACP after considering the following major and minor comments.

- 1- My major concern relates to the discussion of the role of deposition on wet leaf surfaces. From
   the presented material I do not see proof from the model comparison that deposition on wet
   leafs is a significant process. This may be in parts owed to the fact that description and interpre tation are not detailed enough. The authors may address the following questions and issues.
- 27 2- We understand the concern of the reviewer. We use a canopy exchange model which has been previously extensively evaluated for different ecosystems including tropical rainforest. We have 28 29 further stressed this feature of model evaluation and have included some further references for 30 more detailed information about some of the canopy model features relevant to the presented subject of BVOC exchange, e.g., estimating canopy wetness as a function of RH. We are not 31 claiming that deposition to leaf wet surfaces is a significant process. One would actually indeed 32 33 expect this sink to be of minor relevance for the terpenes included in this study given their low solubility's. However, the inferred wet vegetation uptake resistance following the widely-used 34

Wesely (1989) approach of  $\sim 300$  s m<sup>-1</sup> results in simulation of an apparent not ignorable sink 35 also given the large area of potential wet surface. Using this inferred uptake rate we achieved a 36 relative good agreement between model simulated and observed monoterpene mixing ratios, not 37 only in magnitude (which is no surprise given the fact that the constant basal emission factor has 38 been selected to reproduce the campaign average mixing ratios) but especially regarding the di-39 urnal dynamics. The results point to a combined effect of the potential role of chemistry and 40 deposition as monoterpene sinks pointing to the necessity to explore further this deposition to 41 42 wet surface process.

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# 1- (I think a paragraph should be started in line 406 where the leaf wetness is discussed.)

- 45 2- We have separated a paragraph for sink processes.
- 47 1. Did you observe the leaf surface wetness in the two nights? If yes, was it the same or was it different?
- 2- Unfortunately we did not have measurements of leaf surface wetness available for the site, and
  therefore we used relative humidity as a proxy for the calculation of leaf surface wetness. MLCCHEM uses relative humidity as a proxy for the fraction of the wet leaf surface (Lammel, 1999).
  This results in smaller estimates of canopy wetness on 17 October 2015 compared to the following days. In the revised version, we now discuss that these estimates of canopy wetness cannot
  be corroborated without measured canopy wetness.
- You suggest that the model has a strong missing sink in the night from the 17. to the 18. Is this
  the only possible explanation for the strongly overpredicted a-pinene mixing ratio? Or could be
  there more reasons for the model showing so high a-pinene in the night from the 17. to the 18.?
  Only if the source and the chemical sink of a-pinene are about the same in the two nights then
  the depositions sink must have been also similar as the mixing ratios are about the same. Were
  source and chemical sink the same in both nights?
- 2. In response to this comment we modified the discussion about the comparison of the simulated 61 and observed temporal variability in monoterpene mixing ratios and the role of canopy wetness. 62 Rather than stressing these contrasts between the night of 17-18 October and the other nights we 63 now contrast more generally the 17<sup>th</sup> and the other three days. This approach was chosen due to 64 missing meteorological observations for the 18<sup>th</sup>. There is a significant change in the modeled 65 canopy wetness conditions around midnight due to using a prescribed typical diurnal cycle in 66 some meteorological parameters instead of the actual observations for the 17<sup>th</sup>. It provides 67 though an interesting sensitivity analysis since as soon as the imposed canopy wetness increases 68 to a value of 1,  $\alpha$ -pinene drops to much lower values. This strong dependence of the simulated 69 temporal variability on changes in canopy wetness conditions is confirmed by an analysis of the 70 process tendencies in the model which show that the emissions do not significantly change com-71 paring the 17<sup>th</sup> with the other 3 days of the campaign. The contribution by chemical oxidation 72

- 73also shows some temporal variability due to differences in  $O_3$ , OH and  $NO_3$  but this sink is rela-74tively small compared to the dominant sink of  $\alpha$ -pinene, canopy deposition. There the changes in75canopy wetness contrasting the 17<sup>th</sup> with the other 3 days explains to a large extent the simulated76temporal variability in mixing ratios of  $\alpha$ -pinene and the other terpenes and which generally77agrees quite well with the observed temporal variability.
- How can deposition on wet leaves affect so strongly the mixing ratio of non-water soluble compounds like a-pinene?
- 2- What we present here, relies on application of inferred uptake efficiencies for these monoter-80 penes not only commonly being used in any model study following the Wesley (1989) approach 81 to consider dry deposition but also recently being applied in a detailed study on Boreal forest 82 canopy exchange of BVOCs (Zhou et al., 2017). The inferred monoterpene wet-surface uptake 83 resistances on the order of 300 s m<sup>-1</sup>, based on Henry law's constants on the order of  $\sim 1-3^{e-2}$  M 84 atm<sup>-1</sup>, suggest a potentially still quite efficient uptake by leaf wetness. Note that this seems to be 85 similar to the observed removal of ozone by wet canopies where, despite its low solubility, there 86 is actual experimental evidence that canopy wetness enhances O<sub>3</sub> deposition potentially due to 87 aqueous-phase chemical interactions. Clearly, this feature on leaf wetness in monoterpene re-88 moval has to be corroborated by further experimental evidence. 89
  - 1- I understand there was no RH data for the 18. How could the model derive a reliable surface wetness then?
  - 2- See also our response to the previous points about missing canopy wetness observations and the impact of the imposed meteorological data for the 18<sup>th</sup> of October.
- Why do the different MT respond so differently to the ill predicted surface wetness on 17.: a pinene and b-pinene too high, limonene ok, myrcene in between, and terpinolene too low. Inso far I think the statement in line 527 is not justified by the presented material and analysis.
- 97 2- In the revised version we have excluded the respective statement.
- 98 Major comments:

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- line 514-518: I do not understand what the authors are discussing here. I cannot understand how
   a "potential" aerosol growth could have a dimension of molecules per cm3 and an order of 10-5
   -10-6. Moreover, I did not find such notations in Bonn et al. 2014. This part must be clarified or
   deleted.
- 103 2- Our aim was to discuss the effects of mono- and sesquiterpenes as classified in the Amazon on particle formation rates i.e. not aerosol growth rates. Admiringly, the reviewer is correct as the 104 wrong wording was used. The units provided are wrong as the unit needs to be "cm<sup>-3</sup> s<sup>-1</sup>". These 105 have been corrected in the revision. A particle growth would focus on size change per time i.e. 106 nm h<sup>-1</sup> for a certain initial size but not change concentrations per time as provided. In general, 107 we used the word "potential" as no direct measurements of particle growth or of particle size 108 distribution were available and calculations are based on the transfer of results from elsewhere 109 to Brazil. With respect to the aerosol formation and growth rate, it has been shown (for example 110

- by Wolf et al. (2009: 2011) and by Hummel (2010) and included in Bonn et al. (2014)) that the 111 initial growth of (smaller) particles is determined by organic peroxy radicals as one of two pro-112 cesses. Those processes are (a) the reactive interaction of large organic peroxy radicals (RO2) 113 from e.g. mono- and sesquiterpene reactions with OH or NO3 and (b) in case sufficient organic 114 mass has been acquired already uptake by partitioning of oxidized VOCs with a reduced volatil-115 ity. Both processes can be simulated and large organic RO2 concentrations approximated assum-116 ing steady-state conditions for radicals which is a common assumption. For clarification the text 117 has been changed to: 118
- 3- L388 (please refer to supplement for correct location of line number): "Following the equation 119 established by Bonn et al. (2014) (Equation number 5 in text), we were able to estimate the po-120 tential aerosol particle number formation rate initiated by monoterpene species only  $(1 \times 10^{-5} \text{ to})$ 121  $5x10^{-5}$  cm<sup>-3</sup> s<sup>-1</sup> at 24 m) assuming steady state conditions for radicals. Those were found to be 122 approximately two orders of magnitude smaller than the calculated potential new aerosol parti-123 cle formation rate caused by oxidation products of sesquiterpenes. Our calculations assume mix-124 ing ratios of sesquiterpenes of 0.2 ppb revealing potential formation rates of  $1 \times 10^{-3}$  and  $4.5 \times 10^{-3}$ 125 cm<sup>-3</sup> s<sup>-1</sup> at 24 m) based on previous measurements in the Amazon (Jardine et al., 2011) which are 126 remarkably smaller than observed at mid-latitude conditions (Bonn et al., 2014)." 127
- 128
- line 519f: This a strong statement. Do you have proof for that or a reference? Moreover, potential aerosol growth and formation rate are not the same. Indeed NO suppresses nucleation (Wildt et al. ACP, 2014) but not so much the yield (Sarrafzadeh et al. ACP, 2016)
- We kindly refer the reviewer to Sarrafzadeh et al. (2006) and the respective Figure 1 that illustrates how SOA yield depends on the BVOC/NOx ratio (see attached figure 1).
- 134For Amazonian conditions (in contrast to smog chamber measurement conditions) the typical
- BVOC mixing ratios are smaller than 3 ppb and most of the time around 1 ppb. If we apply Fig. 1 of Sarrafzadeh et al. (2016), the situation is located on the very left of the plot. Increasing NOx 1 at constant BVOC mixing ratio will decrease the BVOC/NOx ratio and lead to a decline in SOA 1 yield. The unaffected region of the same figure is not applicable for the Amazon region because 1 of present concentrations. This is in line with the findings of Wildt et al. (2014) for smaller parti-1 cles indicating similar processes responsible for the growth. For clarification we provide with a 1 more detailed explanation as below:
- L396: "Furthermore, the level of present NO (nitric oxide) also affects the potential aerosol growth (Wildt et al., 2014) and yield (Sarrafzadeh et al., 2016) at low BVOC/NOx ratios. As the theory assumes contributions of larger organic peroxy radicals (RO2), which are destroyed by reactions e.g. with NO, increasing NOx at constant BVOC mixing ratio will decrease the BVOC/NOx ratio and lead to a decline in SOA yield. Our calculations showed this effect, with a change of NO from 0.2 ppby to 1 ppby leading to a decrease in the formation rate at a diameter
  - 4

- of 3 nm. This interdependence calls for a consistent consideration of the BVOC and NOx exchange in aerosol formation and growth studies."
- 150 Minor comments:
- 151 1- Starting with line 219 the authors give errors/uncertainty ranges/standard deviations in form of 152 "+/-yx" at many instances (including Table 1) without explaining the specific meaning.
- **153** 2- We have added averages +/- standard deviation where corresponded in the text.
- 154 1- line 156: "142 Td", Td is not explained
- 155 2- This is now explained in the text.
- line 186f: The formulation "selected" suggests to me that the initial EA and the fractional EA for
   individual compounds were used as adaptable parameters. Is this case? If so, then the overall
   good agreement between modelled and observed concentrations in section 3.5 (line 398f) and
   Figure 6 is not surprising. If not, explain the in more detail the rationale for these selections.
- 2- The emission factors from the selected monoterpenes were indeed adaptable parameters. We 160 used selected constant basal emissions for terpenes for the plant ecosystem class, in this case 161 tropical rainforest. We used initially basal emission factor for monoterpenes for tropical rainfor-162 est of 0.4 ug C g<sup>-1</sup> hr<sup>-1</sup> (Guenther et al., 1995). Then we partitioned this emission flux over the 163 different monoterpene species to see how well the model would reproduce the observed mixing 164 ratios. Following, we adjusted the value for each species until the observed 4-day average mix-165 ing ratios were reasonably well reproduced. Given the daily changes in the processes potentially 166 involved in explaining the observed temporal variability, the presented analysis gives you infor-167 mation about the role of the different processes, as these have different dominating effects at dif-168 ferent times of the day and depend, at the same time, on many other drivers. 169
- 170In the text we have chosen to show the basal emissions in mg C  $g^{-1} h^{-1}$ , rather than emission fac-171tor constant, and we have modified the text as follows in order to better explain this issue.
- 3- L179: "The monoterpene basal leaf-scale emission factors have been selected as such that the 172 model simulates monoterpene mixing ratios of comparable magnitude compared to the cam-173 paign average observed mixing ratios. In the evaluation of simulated and observed mixing ratios 174 we mainly focus on the comparison between the simulated and observed temporal variability be-175 ing determined by the differences in canopy processes for contrasting nocturnal and daytime 176 conditions. For the model simulation, the basal emission factors were 0.18  $\mu$ g C g<sup>-1</sup> h<sup>-1</sup> for  $\alpha$ -pi-177 nene, 0.04  $\mu$ g C g<sup>-1</sup> h<sup>-1</sup> for  $\beta$ -pinene, 0.11  $\mu$ g C g<sup>-1</sup> h<sup>-1</sup> for  $\alpha$ -terpinene, 0.9  $\mu$ g C g<sup>-1</sup> h<sup>-1</sup> for limo-178 nene and 0.18  $\mu$ g C g<sup>-1</sup> h<sup>-1</sup> for myrcene. Note the selected relative high basal emission flux for 179 limonene is required to reach simulated mixing ratios comparable to the observed ones." 180
- 1- line 386f: Similar is true for the formulation about the "assumed" lead area index and leaf area density. What is the rationale for these choices? Did you estimate it from observations on the site?
- LAI was previously measured at the site, and we have included this reference in the methodol ogy.

- line 256, first §: These percentages are averages over day and night hours and three measurement days? They should have standard deviations.
- 188 2- We agree with the reviewer and have added the standard deviations of the percentages both in189 the graph and text.
- line 286, first §: I suggest to define what you mean by reactivity. e.g. OH reactivity = kOH\*[MT].
- **192** 2- We have added this clarification in the text.
- 1- line 303, Table 2: The header of the third column is misleading. It should read "normalized reac tivity for 1 ppb [s-1]", or so. Otherwise confusion with the use of reactivity later in the text.
- **195** 2- We have modified this in Table 2 accordingly.
- 196 1- line 489: Is Hallquist 2009 (a review) a good reference for this statement?
- 197 2- We agree this is not the most appropriate reference and we have excluded it.
- line 542f: That is not new and with the given formulation the conclusion in its generality does
   not make sense. The authors did not show that there is no suited representation of MT, which
   considers also their reactivity. Moreover, the degree of tolerable simplification depends also on
   the purpose of the model calculation.
- 202 2- We have reformulated the sentence, and we believe the point of the sentence was not well stated
  203 before.
- 1- L502 "Furthermore, reactivity calculations demonstrated that higher abundance of MT does not 204 automatically imply higher reactivity as the most abundant compounds may not be the most at-205 206 mospheric chemically relevant compound or the relative contribution of different monoterpenes may change. Our calculations support the view to that the role of canopy exchange may be erro-207 neously estimated when not taking into account speciation based reactivity in models. Moreover, 208 simulations with a canopy exchange modelling system to assess the role of canopy interactions 209 compared relatively well with the observed temporal variability in speciated monoterpenes, but 210 also indicate the necessity of more experiments to enhance our understanding of in-canopy sinks 211 of these monoterpenes." 212
- 213 Reviewer #2 (Remarks to the Author):
- 214 General comments:
- 1- The manuscript "Monoterpene chemical speciation in the Amazon tropical rainforest: variation with season, height, and time of day at the Amazon Tall Tower Observatory (ATTO)" is a suitable and in the scope of "Atmosphere Chemistry and Physics", shading light on the importance of monoterpene characterisation. Although, the manuscript is interesting it requires a major revision as it lacks the clear structure. The reader is often forced to go forward and beck. Many discussions are found in Results (e.g. L230-232, L241-242, L267-270, L276-278, L280-281 etc.).
- 221 Many statements are started with not suitable paragraph context (e.g. L306: "The most abundant

- species are  $\alpha$ -pinene, limonene and myrcene".). Some statements and discussion are out of con-
- text or even incorrect as mentioned in the Specific comments below. It is also too long for the scientific content. I suggest the following: Restructure the text in order to join Results and Discussion into one section with adequate subsections, redo the figures as suggested below and de-
- 226 lete the repetitions in the text. Please also find specific and technical comments.
- 2- We thank the reviewer for the fruitful feedback. As suggested, we have now merged the results
   and discussion sections, removed repetitive sentences and redid the figures.
- 229 Specific comments:
- L60-63. Not entirely true. PTR-MS could have a time resolution <1s (not 30 s). It is also recently coupled with FastGC to characterise monoterpenes in < 2min. 1h GCFID is a bit too much for "current" method. Please update the references for all this.</li>
- 233 2- We agree with the reviewer and therefore we have removed that part from the text.
- L103 and 125 Not indicated sampling frequency, 30 min mentioned in L125, is not in agree ment with the resolution presented in the Figure 1b and c. A detailed but structured description
   of the methods used is needed.
- 2- The sampling frequency was every hour for 30 minutes. We agree with the reviewers this is not
   clear in the text and we have modified this paragraph for better explanation.
- 239 3- L111 "The samples were collected from 17 to 20 October 2015 at an hourly frequency. Samples
  240 were collected for 30 min every hour at a flow of 200 cm<sup>3</sup> min<sup>-1</sup> (STP), leading to a collection of
  241 6 L of air in each cartridge."
- L106 Air sampling section needs structure improvement. You first open with your sampler
   giving the reference, then you describe your sampling dates/times and then you come back to the
   sampler, and again back to the sampling procedure.
- 2- We thank the reviewer for this comment, as it is true that section needed improvements. There-fore, we have restructured this section following:
- 3- L97: "Collection of ambient air samples on adsorbent tubes, for subsequent analysis by Gas 247 248 Chromatography – Flame Ionization Detector (GC-FID), was made with two automated cartridge samplers described earlier (Kesselmeier et al., 2002; Kuhn et al., 2002, 2005) positioned 249 at 12 and 24 m on the INSTANT tower. The samplers consist of two main units, a cartridge 250 magazine that holds the adsorbent-filled tubes and the control unit timing the process and re-251 cording the data. This latter unit also houses the pumps (Type N86KT, KNF Neuberger, Frei-252 burg, Germany), pressure gauges, mass flow controllers and power supply. The cartridge maga-253 zine is equipped with solenoid valves controlling the inlet and outlet of up to 20 individual sam-254 255 pling adsorbent tubes. The system is a constant-flow device, with one cartridge position per loop used as a bypass for purging the system. Due to the compact weatherproof housings and the low 256 power consumption, we were able to position one sampler at 24 m and the other one at 12 m, 257 attached to the INSTANT tower booms with commercially available 50 mm aluminium clamps. 258 The adsorbent tubes used for VOC sampling were filled with 130 mg of Carbograph 1 (90 m<sup>2</sup> g<sup>-</sup> 259

- <sup>1</sup>) followed by 130 mg of Carbograph 5 (560 m<sup>2</sup> g<sup>-1</sup>) sorbents. The size of the Carbograph parti-260 cles was in the range of 20–40 mesh. Carbographs 1 and 5 were provided by L.A.R.A s.r.l. 261 (Rome, Italy) (Kesselmeier et al., 2002). The samples were collected from 17 to 20 October 262 2015. Samples were taken for 30 min every hour at a flow of 200 cm<sup>3</sup> min<sup>-1</sup> (STP), leading to a 263 collection of 6 L of air in each cartridge using the automatic sampler. Additional sampling was 264 performed at 24 m with a GSA SG-10-2 personal sampler pump during the years 2012-2014. 265 These earlier samples were collected in the same type of adsorbent tubes as for the automatic 266 sampler, and were filled at 167 cm<sup>3</sup> min<sup>-1</sup> (STP) air flow for 20 min. These additional measure-267 ments took place on 19 and 28 November 2012; 1, 3 and 4 March 2013; 11 to 14 June 2013; 22, 268 25 and 26 September 2013 and on 17 and 21 August 2014." 269
- 1-L224 In Table 1 it is not clear what the tolerance is.
- 271 2- We have added this information to the table legend.
- L128 The "Instrument for chemical analysis". Are you intent to describe the instruments or the methods used?
- 2- We intent to describe the methods used, therefore we have changed the name of this section to
  "Instruments used for chemical analysis".
- 276 1- L137 What is "rapidly" in this context?
- 277 3. We wanted to say that the temperature increased at a high rate. We have removed this term from the text.
- 1- L160 Why just 12 and 24 if all the data are available?
- 2- We only wanted to use 12 and 24m data of the PTR-MS as these were the only heights were
   GC-FID sampling took place in parallel.
- L164 "2σ of the background" 3σ it is more acceptable. Also briefly describe the blanks in
   each sampling systems.
- 2- We have modified the LOD to 3σ. A more detailed explanation of the blanks has been added as following:
- 286 3- L148: "Hourly background measurements with a catalytic converter (Supelco, Inc. with platinum pellets heated to >400°C) and weekly humidity dependent calibrations of the PTR-MS were performed"
- L153: "The compounds measured were monoterpenes (m/z 137) and isoprene (m/z 69). The
  limit of detection (LOD) of the PTR-MS for total monoterpenes was 0.1 ppb and 0.2 ppb for isoprene, determined as 3σ of the background noise."
- 1- L239 Isoprene! Not mentions in material and methods. How is this measured?
- 2- We understand we missed this part in the materials and methods. Isoprene was measured by the
   PTR-MS and by GC-FID. We have added this to the text.
- L253 Figure not clear. E.g. isoprene 24m not visible after 12:00. Legend not descriptive, not
   explained what are the error bars. Some error bars below zero and yet, above LOD? Explain?
- 297 2- This figure has been modified (see attached figure 2)

- Now isoprene mixing ratios as measured by the GC-FID can be better seen. Furthermore, and particularly for the low mixing ratio compounds, it is possible that the measurement is 0, as it can be seen for a-terpenene, but it is not always 0. The data expressed here is an hourly average over three days and samples were collected every hour for 3 days, so having this in mind, standard deviations may be high. On the other hand, LOD for the GC-FID system is around 0.2 ppt, and this is generally way below than measured mixing ratios of the monoterpene species.
- 304 3- L737: "Figure 2: Average diel cycles for α-pinene (a), limonene (b), myrcene (c), ρ-cymene (d),
   305 β-pinene (e) and α-terpinene (f) mixing ratios for 24 m (dashed line) and 12 m (thick line). In the
   306 back, average diel cycle of isoprene mixing ratios as measured by the GC-FID are shown for 24
   307 m (light green) and 24 m (dark green). Error bars represent the standard deviation of the averages."
- 309
- L345- "..in previous campaigns.. earlier samples were collected using a GSA SG10-2 personal
   pump sampler. Adsorbent tubes were filled at 167 cm3 min-1 (STP) air flow for 20 min". This is
   again an example of poor structure. This need to be in the Material and Methods section.
- 313 2- We agree with the reviewer this explanation was not well structured. We have put this information in the methods paper and have removed this from the results.
- 315 3- L113: "Furthermore, additional sampling was performed at 24 m with a GSA SG-10-2 personal sampler pump during the years 2012-2014. These earlier samples were collected in the same type of adsorbent tubes as for the automatic sampler, and were filled at 167 cm3 min-1 (STP) air flow for 20 min. These additional measurements took place on 19 and 28 November 2012; 1, 3 and 4 March 2013; 11 to 14 June 2013; 22, 25 and 26 September 2013 and on 17 and 21 August 2014."
- L380 How the above-canopy O3 concentrations used in the model and for the reactivity calculation represent real situation in the canopy (between 12 and 24m)?
- We have used measured ozone mixing ratios at 12 and 24m for our calculations and not the
   above canopy ozone mixing ratios. We have removed this part from the results as it was already
   specified in the methodology and we clarified that the ozone mixing ratio levels used for reactiv ity calculations were at 12 and 24m.
- 327 The revised text reads as:
- 3- L191: "The simulations with MLC-CHEM were constrained with the observed surface layer net radiation (above the canopy only), wind speed, relative humidity and O<sub>3</sub> mixing ratios as well as the temperature measured above and inside the canopy (8 different heights including 12 and 24m) from 17 to 20 October 2015, coinciding with the measurement dates."
- 332 L310: "For ozone reactivity calculations, 12 ppb was used, as this mixing ratio was observed
- during the measurement period. NO<sub>3</sub> mixing ratios were taken from the MLC-CHEM model
- simulations that predicted mixing ratios of ~0.4 ppt."

- L439 and on "The emission of monoterpenes has generally been thought to be from storage glands in specialized structures like resin ducts, glandular trichomes or related structures (Schürmann et al., 1993; Steinbrecher, 1989)." This generalization and discussion based further in the text are related only to conifer type or monoterpene emitters (see also in your references). Thus, this is irrelevant (and incomparable) for the tropical forest as the physiology (and chemotypes) for monoterpene emission in broadleaf tree species is different to the pool emitters.
- 341 2- We agree with the reviewer that this is an unspecific generalization and we have revised the text342 accordingly:
- 3- L260: "In contrast to plant species of cooler climates, such as spruce, which emit terpenes from 343 pools (Ghirardo et al., 2010; Lerdau et al., 1997), Amazonian plant species have been found to 344 show an emis-sion dependency on light and temperature (Bracho-Nunez et al., 2013; Jardine et 345 al., 2015; Kuhn et al., 2002, 2004). This could partly explain the diurnal pattern of  $\alpha$ -pinene 346 mixing ratios, which exhibit some relation to a light and temperature dependent emission flux 347 (Kuhn et al., 2002; Rinne et al., 2002; Williams et al., 2007). However, this behaviour is not ob-348 served for all monoterpene species. Therefore, the observed diurnal cycles of some monoterpene 349 350 species might be related to a stronger temperature response. While monoterpenes are stored in leaves and their release from these pools is governed by leaf temperature (Monson et al., 1995). 351 Amazonian plant species have been found to show an emission de-pendency on light and tem-352 perature (Bracho-Nunez et al., 2013; Jardine et al., 2015; Kuhn et al., 2002, 2004). This could 353 partly explain the diurnal pattern of  $\alpha$ -pinene mixing ratios, which exhibit some relation to a 354 355 light and temperature dependent emission flux (Kuhn et al., 2002; Rinne et al., 2002; Williams et al., 2007). However, this behaviour is not observed for all monoterpene species. Therefore, 356 357 the observed diurnal cycles of some monoterpene species might be triggered by stronger temperature dependencies." 358
- 1- L502 and 522 Any leaf level experiments to address this? Needs a brief discussion.
- We assume that the reviewer is referring to L520-522. Unfortunately, we do not have any leaf
   level experiments to address this. We have performed these calculations which give an estima tion of the possible role of NO in SOA growth indicating the need for further studies. We only
   want to point out the necessity of assessing this role when studying SOA growth from monoter pene species in the Amazon region.
- 365 Technical corrections:
- 366 1- L1 Two words "Amazon" in a title are not needed.
- **367** 2- We have removed Amazon from the title.
- 368 1- L19 Why just in Amazon rainforest? You may just say "a rainforest".
- **369 2-** We have changed it accordingly.
- 1- L24 "automatic" automated?
- **371** 2- We have changed it accordingly.

372	1-	L31 – "may not be the most atmospheric chemically relevant compounds". Although it might be						
3/3	2	We have changed this sentence to: "Peactivity calculations showed that higher abundance does						
374	2-	not imply higher reactivity?						
375	1	174.76 "approx" a bit over repetition of this word throughout the manuscript and here						
376	1-	L/4-/0 – scarce – a bit over-repetition of this word throughout the manuscript and here.						
3//	Z- 1	L 220 The is a discussion in the "Desult" section						
378	1-	L250 – The is a discussion in the Result section.						
379	Z- 1	L 252 Element 2 les endenderes tert ter energil						
380	1-	L255 – Figure 2 legend and axes text too small.						
381	Z- 1	We have modified this figure and increased legend and axes.						
382	1-	L283 – Figure 3 – figure caption above not needed. Description needs to be extended.						
383	2-	we have modified this figure and elaborated further in the description:						
384	3-	"Pie charts representing the day (a and c) and night (b and d)day and night averaged monoter-						
385		pene species abundance in aver-age percentage with standard deviation at 24 (a and b) and 24 (c						
386		and d) m. Day period was from 0900h to 1700h and night period was from 2000h to 0500h."						
387								
388	1-	L289 – "O1D" ?!						
389	2-	We agree with the reviewer this was not specified in the text and we have removed it.						
390	1-	L312 – Figure 4 out of margins, figure caption not needed. Use a) b): : : to refer to the individual						
391		figures with a clear description. Also, check all the figures to meet this standard.						
392	2-	We agree with the reviewer and we have modified the graph and footnote accordingly.						
393	1-	L381-384 - Repetition						
394	2-	We have deleted this part in the results sections to avoid repetition.						
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395	Ref.:							
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- 413
- Monoterpene chemical speciation in <u>the Amazona</u> tropical rainforest: variation with season, height,
   and time of day at the Amazon Tall Tower Observatory (ATTO)
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- 430 431 432
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- 433 Speciated monoterpene measurements in the Amazon rainforest air are scarce, but they are important in order to understand their essential for understanding the contribution of these compounds to the overall 434 reactivity of volatile organic compound (VOCs) emissions towards the main atmospheric oxidants, such 435 as hydroxyl radical (OH), ozone (O<sub>3</sub>) and nitrate radical (NO<sub>3</sub>). In this study, we present the chemical 436 437 speciation of gas phase monoterpenes measured in the tropical rainforest at the Amazon Tall Tower Ob-438 servatory (ATTO, Amazonas, Brazil). Samples of VOCs were collected by two automated sampling systems positioned on a tower at 12 and 24 m height and analysed using Gas Chromatography Flame 439 Ionization Detection (GC-FID). The samples were collected in October 2015, representing the dry sea-440 son, and compared with previous wet and dry season studies at the site. In addition, vertical profile 441 442 measurements (at 12 and 24 m) of total monoterpene mixing ratios were made using Proton-Transfer 443 Reaction Mass Spectrometry (PTR-MS). The results showed a distinctly different chemical speciation 444 between day and night. For instance,  $\alpha$ -pinene was more abundant during the day, whereas limonene 445 was more abundant at night. Reactivity calculations showed that higher abundance does not generally imply higher reactivity. Furthermore, inter- and intra-annual results demonstrate similar chemodiversity 446 during the dry seasons analysed. Simulations with a canopy exchange modelling system show simulated 447

monoterpene mixing ratios that compare relatively well with the observed mixing ratios, but also indi cate the necessity of more experiments to enhance our understanding of in-canopy sinks of these mono terpenescompounds.

451

### 452 **1. Introduction**

453 Isoprenoids such as isoprene ( $C_5H_8$ ), monoterpenes ( $C_{10}H_{16}$ ) and sesquiterpenes ( $C_{15}H_{24}$ ) are 454 considered to be key contributors to the production of biogenic secondary organic aerosol (SOA), which 455 affects cloud condensation nuclei production (Engelhart et al., 2008; Jokinen et al., 2015; Pöschl et al., 456 2010). While isoprene has been shown to be a globally significant source of SOA (Claeves et al., 457 2004), it's presence has been also shown to can also inhibit SOA formation under certain conditions 458 (Kiendler-Scharr et al., 2009). By virtue of their lower volatility and higher ozone reactivity, monoter-459 penes and sesquiterpenes are strong sources of secondary organic aerosol (SOA) through the generation 460 of low-volatility oxidation products formed via ozonolysis and hydroxyl radical oxidation (Bonn and 461 Moortgat, 2003; Zhao et al., 2015).

462

463 The main source of monoterpenes to the global atmosphere is emission from vegetation, with 464 smaller contributions from soil (Kesselmeier and Staudt, 1999; Kuhn et al., 2002; Ormeno et al., 2007). Synthesis of the monoterpene species occurs via the non-mevalonate pathway within the plant chloro-465 466 plast (Kesselmeier and Staudt, 1999; Lichtenthaler, 1999; Schwender et al., 1996), which explains the 467 light dependency also known to determine isoprene synthesis and emission. These commonly emitted 468 compounds have been identified as important signalling compounds through plant-to-plant, plant-insect or plant-microbe interactions (Gershenzon, 2007; Gershenzon and Dudareva, 2007; Kishimoto et al., 469 470 2006; Maag et al., 2015) and they are thought to protect photosynthetic membranes against abiotic 471 stresses (Jardine et al., 2017; Penuelas and Llusia, 2002; Vickers et al., 2009).

472

473 Despite having a common sum formula, variations in the molecular structure of the various
474 monoterpenes result in large variations (over two orders of magnitude) of their reaction rate coefficients
475 with the hydroxyl radical (OH), ozone (O<sub>3</sub>) and nitrate radical (NO<sub>3</sub>). Current quantification methods

476 for monoterpenes include the fast (ca. 30 seconds) but unspecific total monoterpene measurement by the proton transfer reaction mass spectrometer (PTR-MS) (Lindinger and Jordan, 1998) and the slower (ca. 477 1 hour) but chemically speciated gas chromatographic methods. In order to gauge the role of these spe-478 479 cies in atmospheric chemistry, chemical speciation needs to be provided by the gas chromatographic techniques while the on-line mass spectrometer can assess how the total mass changes over time. It 480 481 should be noted that the structure and reactivity of the monoterpenes can also have This leads to differ-482 ent implications for the efficiency of SOA formation (Hallquist et al., 2009; Kiendler-Scharr et al., 483 2009; Mentel et al., 2009; O'Dowd et al., 2002). In most cases, SOA products are poorly characterized due to a scarcity of measurements (Martin et al., 2010). 484

485

Considering the overall size of the Amazon rainforest (5.4 million  $\text{km}^2$  in 2001; Malhi et al., 486 2008) and the significant contribution of BVOC emissions from this vast forest to the global VOC 487 budget (globally 1000 Tg of carbon yr<sup>-1</sup>; Guenther et al., 2012), measurements of total monoterpene 488 489 emissions and mixing ratios from this ecosystem are scarce (Greenberg and Zimmerman, 1984; Helmig 490 et al., 1998; Jardine et al., 2015, 2011, 2017; Karl et al., 2007; Rinne et al., 2002; Yáñez-Serrano et al., 491 2015). Speciated measurements are even more rare (Jardine et al., 2015, 2017; Kesselmeier et al., 2002; 492 Kuhn et al., 2004). Yet, this information is essential for our understanding of the functioning of the Amazon rainforest in atmospheric chemistry-climate interactions. Knowledge of these processes also serves 493 494 to improve predictions of future changes in atmospheric composition and to assess the impact of 495 changes in regional emissions and land use on global climate caused by Amazon deforestation.

496

In this study, we evaluate measurements of speciated rainforest monoterpene mixing ratios as a function of height in the canopy, season and diel cycle. This evaluation includes a comparison with a canopy exchange modelling system (MLC-CHEM, Multi-Layer Canopy Chemistry Exchange model) to support analysis of the measured temporal variability in speciated rainforest monoterpene mixing ratios <u>inside the tropical rainforest canopy. The MLC-CHEM model</u> was also <u>selected</u> since it has been already extensively applied for site- to global-scale studies on atmosphere-biosphere exchange for tropical rainforests ecosystems (Ganzeveld et al., 2008, 2002; Ganzeveld and Lelieveld, 2004; Kuhn et al.,
2010).

505

### 506 2. Methodology

507 2.1. Site

508 The site chosen for this study was the Amazon Tall Tower Observatory, ATTO (Andreae et al., 2015). This site is located in Central Amazonia (S 02° 08.647' W 58° 59.992'), 150 km NE of the clos-509 510 est <del>populated</del> large city, Manaus, Brazil, Due to the prevailing north-easterly wind direction, the influence of the Manaus plume is negligible and the measurements at this site can be considered to reflect 511 512 pristine tropical forest conditions affected by air masses that have passed over about 1000 km of undis-513 turbed rainforest. The site is equipped with a 325 m tall tower as well as two smaller towers. This study was carried out on the INSTANT tower, an 80-m walk-up tower located 600 m from the tall tower in 514 515 easterly direction. Sampling was performed on this tower below canopy top (mean canopy height 35 m) 516 at two different heights (12 m and 24 m). For a comprehensive site description see Andreae et al. (2015). 517

518

519

2.2. Air sampling

520 Collection of ambient air samples on adsorbent tubes, for subsequent analysis by Gas Chroma-521 tography – Flame Ionization Detector (GC-FID), was made with two automated cartridge samplers de-522 scribed earlier (Kesselmeier et al., 2002; Kuhn et al., 2002, 2005) positioned at 12 and 24 m on the IN-523 STANT tower. The samplers consist of two main units, a cartridge magazine that holds the adsorbent-524 filled tubes and the control unit timing the process and recording the data. This latter unit also houses the pumps (Type N86KT, KNF Neuberger, Freiburg, Germany), pressure gauges, mass flow controllers 525 and power supply. The cartridge magazine is equipped with solenoid valves controlling the inlet and 526 527 outlet of up to 20 individual sampling adsorbent tubes. The system is a constant-flow device, with one 528 cartridge position per loop used as a bypass for purging the system. Due to the compact weatherproof

529 housings and the low power consumption, we were able to position one sampler at 24 m and the other one at 12 m, attached to the INSTANT tower booms with commercially available 50 mm aluminium 530 531 clamps. The height of the canopy is approximately 35 m (Andreae et al., 2015). The adsorbent tubes used for VOC sampling were filled with 130 mg of Carbograph 1 (90 m<sup>2</sup> g<sup>-1</sup>) followed by 130 mg of 532 Carbograph 5 (560 m<sup>2</sup> g<sup>-1</sup>) sorbents. The size of the Carbograph particles was in the range of 20–40 533 mesh. Carbographs 1 and 5 were provided by L.A.R.A s.r.l. (Rome, Italy) (Kesselmeier et al., 2002). 534 535 The samples were collected from 17 to 20 October 2015. Samples were taken for 30 min every hour at a flow of 200 cm<sup>3</sup> min<sup>-1</sup> (STP), leading to a collection of 6 L of air in each cartridge using the automatic 536 537 sampler. Additional sampling was performed at 24 m with a GSA SG-10-2 personal sampler pump dur-538 ing the years 2012-2014. These earlier samples were collected in the same type of adsorbent tubes as for the automatic sampler, and were filled at 167 cm<sup>3</sup> min<sup>-1</sup> (STP) air flow for 20 min. These additional 539 measurements took place on 19 and 28 November 2012; 1, 3 and 4 March 2013; 11 to 14 June 2013; 22, 540 541 25 and 26 September 2013 and on 17 and 21 August 2014.

- 542
- 543

2.3. Instruments used for chemical analysis

544 2.3.1. Gas Chromatography –Flame Ionization Detector (GC-FID)

545 After collection, the adsorbent tubes were analysed at the Max Planck Institute for Chemistry (MPIC) using a Gas Chromatograph equipped with a Flame Ionization Detector (GC-FID, Model Auto-546 547 System XL, Perkin Elmer GmbH, Germany) for identification and quantification of the monoterpene 548 species. Helium was used as carrier gas, and -Sseparation occurred on a 100 meter HP-1 column with 549 0.22 mm inner diameter, and coated with the non-polar dimethylpolysiloxane as stationary phase. The 550 compound mixture collected in the adsorbent tubes was discharged into the gas stream with the help of 551 a two-step desorption system (Model ATD400, Perkin Elmer, Germany). The samples were cryofocused in a cold trap at -30 °C filled with Carbograph 5, providing better defined peaks in the chromato-552 553 grams. Afterwards the cold trap was rapidly heated to 280°C and the pre-concentrated sample injected onto the column. The following temperature programme was used: (-10 to 40 °C at 20 °C min<sup>-1</sup>, 40 to 554 145 °C at 1.5 °C min<sup>-1</sup>, and 145 to 220 °C at 30 °C min<sup>-1</sup>). The separated compounds were quantified 555

with a Flame Ionization Detector (FID). Identification was achieved through spiked injection of pure
compounds. For a more detailed description see Kesselmeier et al., (2002).

559	Calibration for VOCs containing no heteroatoms was achieved by using a standard gas mixture
560	of isoprene and several n-alkanes (n-pentane, n-hexane, n-heptane, n-octane, n-nonane, and n-decane)
561	(Apel-Riemer Environmental Inc., USA). In this case, it is assumed that the "effective carbon number"
562	(Sternberg et al., 1962) is equal to the real carbon number of the molecules (Komenda, 2001), yielding a
563	signal response that is proportional to the real carbon number. The monoterpenes identified and quanti-
564	fied were $\alpha$ -pinene, camphene, sabinene, $\beta$ -pinene, myrcene, $\alpha$ -phellandrene, 3-carene, $\alpha$ -terpinene, $\rho$ -
565	cymene, limonene and $\gamma$ -terpinene. Isoprene was also quantified. The detection limit for the GC-FID
566	was 2 ppt (Bracho-Nunez et al., 2011).
567	
568	2.3.2. Proton Transfer Reaction - Mass Spectrometer (PTR-MS)
569	Online total monoterpene mixing ratios were determined by a quadrupole Proton Transfer Reac-
570	tion - Mass Spectrometer, PTR-MS (Ionicon Analytic, Austria). The PTR-MS was operated under
571	standard conditions (2.2 mbar drift pressure, 600 V drift voltage, with an E/N of 142 Townsend (Td)).
572	In addition to weekly humidity dependent calibrations, hourly background measurements were per-
573	formed with a catalytic converter (Supelco, Inc. with platinum pellets heated to >400°C). A gravimetri-
574	cally prepared multicomponent standard for calibration was obtained from Apel & Riemer Environmen-
575	tal, USA. The measurements were carried out at two different heights (0.05, 0.5, 4, 12 and, 24, 53 and
576	79 m) with the PTR-MS switching sequentially between each height at 2 min intervals-and only data
577	from 12 and 24 m is shown. The inlet lines were made of PTFE (9.5 mm OD), insulated and heated to
578	50 °C, and had PTFE particle inlet filters at the intake end. The compounds of interest for this study
579	were the isoprene ( $m/z$ 69) and the sum of monoterpenes ( $m/z$ 137) and isoprene (in $m/z$ 69). The limit
•	

of detection (LOD) of the PTR-MS for total monoterpenes was 0.1 ppb<u>and</u> 0.2 ppb <u>for isoprene</u>, determined as 3σ of the background noise. More information about the gradient system and PTR-MS operation at ATTO can be found elsewhere (Nölscher et al., 2016; Yáñez-Serrano et al., 2015).

583 584

### 2.4. Multi-Layer Canopy Chemistry Exchange model (MLC-CHEM)

585 To analyse the magnitude and temporal variability inof the observed monoterpene concentrations inside and above the forest canopy, we applied the Multi-Laver Canopy Chemistry Exchange 586 Model (MLC-CHEM), driven by the observed micro-meteorology and ozone surface layer mixing ra-587 tios. MLC-CHEM was originally developed and implemented in a single-column model (SCM). It is 588 originally set up as well as also in a global chemistry and climate-modelling system to assess the role of 589 590 canopy processes in local- to global-scale atmosphere-biosphere exchange of nitrogen oxides (Ganzeveld et al., 2008, 2002; Kuhn et al., 2010). The model's generalized representation of chemistry, 591 dry deposition, emissions and turbulent mixing allows to studying the role of canopy interactions in de-592 termining atmosphere-biosphere exchange fluxes and in-canopy and surface layer mixing ratios of, e.g., 593 ozone  $(O_3)$ , nitrogen oxides  $(NO_x)$  and biogenic volatile organic compounds (BVOCs). The BVOC 594 emissions are calculated according to MEGAN (Guenther et al., 2006), considering the vertical distribu-595 596 tion of biomass and direct as well as diffuse radiation to calculate leaf-scale BVOC emissions. The current implementation of canopy chemistry in MLC-CHEM considers, in addition to standard photo--597 598 chemistry involving  $O_3$ ,  $NO_x$ , methane (CH<sub>4</sub>) and carbon monoxide (CO), the role of non-methane hydrocarbons including isoprene, and a selection of hydrocarbon oxidation products such as formalde-599 hyde, higher aldehydes and acetone. Oxidation of the monoterpenes by OH,  $O_3$  and  $NO_3$  is taken into 600 601 account, but the role of the monoterpene oxidation products in photo-chemistry is not considered in the current implementation of the chemistry scheme in MLC-CHEM. For this study, we have extended 602 MLC-CHEM to consider, besides the already included compounds  $\alpha$ -pinene and  $\beta$ -pinene, the observed 603 monoterpene species,  $\alpha$ -terpinene, limonene and myrcene. The monoterpene basal leaf-scale monoter-604 605 pene emission factors have been selected such that the model simulates monoterpene mixing ratios of comparable magnitude compared to the campaign-average observed mixing ratios. In the evaluation of 606

- 607 simulated and observed mixing ratios we mainly focus on comparison of the simulated and observed 608 temporal variability being determined by the differences in canopy processes for contrasting nocturnal and daytime conditions. For the model simulation, the basal emission factors were 0.18  $\mu$ g C g<sup>-1</sup> h<sup>-1</sup> for 609  $\alpha$ -pinene, 0.04 ug C g<sup>-1</sup> h<sup>-1</sup> for  $\beta$ -pinene, 0.11 ug C g<sup>-1</sup> h<sup>-1</sup> for  $\alpha$ -terpinene, 0.9 ug C g<sup>-1</sup> h<sup>-1</sup> for limonene 610 and 0.18 ug C  $g^{-1}$  h<sup>-1</sup> for myrcene. Note the selected relative high basal emission flux for limonene is 611 required to arrive at simulated mixing ratios comparable to the observed ones. Regarding the physical 612 sinks: dry deposition of gases including the BVOC compounds depends on their uptake resistances cal-613 614 culated according to Weselv's (1989) parameterization, which estimates these uptake resistances based on the compounds' solubility and reactivity. 615
- 616

617 The simulations with MLC-CHEM were constrained with the observed surface layer net radiation (above the canopy only), wind speed, relative humidity and  $O_3$  mixing ratios as well as the temper-618 atures measured above and inside the canopy (8 different heights including 12 and 24m) from 17 to 20 619 October 2015, coinciding with the measurement dates. These simulations represent a set-up of MLC-620 CHEM distinguishing six canopy levels with a canopy height of 30 m, implying canopy layers with a 621 thickness of 5 m. Furthermore, we assumed a Leaf Area Index (LAI) of 5 m<sup>2</sup> m<sup>-2</sup> and a Leaf Area Den-622 sity (LAD) profile such that about 70% of this biomass is present in the top 15 m of the canopy, as pre-623 624 viously observed in tropical rainforest sites (Nölscher et al., 2016). Monoterpene emissions by 625 vegetation were simulated using a temperature-only dependent emission flux as a function of the amount of biomass in each layer and the measured canopy temperature profiles interpolating between 626 the 0.4 m and 26 m temperature sensors. Meteorological observations for 18 October were missing and 627 therefore MLC-CHEM was constrained for this day by first-order estimates of the diurnal cycles in radi-628 629 ation, air and surface temperatures, relative humidity and wind speed comparable to the previous and posteriorsubsequent day's meteorological conditions. 630

631

# 632 **3. Results and discussion**

- 633
- 3.1. Time series and diel cycles

634 The continuous online PTR-MS measurements were compared with off-line GC-FID samples over the course of 3 days in October 2015 (Figure 1). The close agreement between the two measure-635 636 ment techniques provides confidence that almost all monoterpenes present at in ambient air in at the site were being measured. Note that in this comparison, o-cymene (an aromatic monoterpene) was removed 637 from the calculations as the PTR-MS does not detect it on m/z 137. The observed differences in the 638 monoterpene chemodiversity at in the ATTO-rainforest canopy atmosphere were regarded to be driven 639 by differences in emission, reactivity withto the oxidizing species, physical removal processes and tur-640 641 bulent mixing conditions.

642

643 The total monoterpene mixing ratios were higher during the day, when temperature and solar 644 radiation were at their maxima. Most of the observed distinct diurnal cycle in total monoterpene mixing 645 ratios couldan be attributed to  $\alpha$ -pinene, which was the dominant species during the daytime (0900h to 646 1700h) with mixing ratios as large as (average  $\pm$  standard deviation) 0.33 $\pm$ 0.04 and 0.38 $\pm$ 0.21 ppb at 12 647 and 24 m respectively, and  $0.15\pm0.05$  and  $0.11\pm0.06$  ppb for the night (2000h to 0500h) at 12 and 24 m. The second most abundant monoterpene species was limonene, with observed average daytime mixing 648 649 ratios of  $0.18\pm0.09$  and  $0.19\pm0.12$  ppb at 12 and 24 m, respectively, and  $0.18\pm0.01$  and  $0.14\pm0.07$  ppb 650 for the night time at 12 and 24 m.

651

652 When comparing our results to previously published studies, we observed consistent differences with other regions of the Amazon rainforest. For instance, Kesselmeier et al. (2002), studied the sea-653 sonal monoterpene speciation in the Rondonia rainforest in southern Amazonia. Even though they 654 found the same monoterpene species as presented in this study, their individual abundances were very 655 656 different compared to the mixing ratios for the dry season at the ATTO site.  $\alpha$ -pipinenene and limonene were much higher at ATTO than in Rondonia, whereas camphene was substantially lower. In the case 657 of β-pinene, the abundance measured at ATTO was much lower than at other Amazonian sites (Andreae 658 et al., 2002; Karl et al., 2007). DespiteGiven that emission patterns are highly dependent on species, en-659 vironmental conditions and stresses, these differences show underline that it cannot be assumed that the 660 same speciation and emission rates of monoterpenes exist throughout the vast Amazon basin. 661

668

<u>Furthermore, the difference between the 12 and 24 m height total monoterpene mixing ratios</u>
 <u>wasis</u> minor given the variance of the measurements, but there <u>wasis</u> a tendency for the difference to be
 more pronounced during night time (Table 1). <u>These more pronounced differences between the meas-</u>
 <u>urement heights could be also due to an enhanced sensitivity of nocturnal mixing ratios to small</u>
 <u>changes in source and sink terms for the suppressed mixing conditions prevailing during the night time.</u>

669 The continuous online measurements by the quadrupole PTR-MS indicated a clear diurnal cycle in the measured mixing ratios of the sum of monoterpenes, which has been reported previously from 670 671 this site (Yáñez-Serrano et al., 2015). In order to assess the effect of each individual monoterpene species, we further investigated their diurnal cycles as obtained by the off-line GC-FID samples. The meas-672 ured diel cycles for the most relevant monoterpene species at the ATTO site were very similar at both 673 heights. We also compared the measured diel cycles of isoprene as measured by the GC-FID with the 674 675 observed diel cycle for the different monoterpene species for 12 and 24 m. The compounds that showed 676 a diurnal cycle similar to isoprene were  $\alpha$ -pinene and  $\rho$ -cymene (Figure 2). This could be due to the 677 emission of  $\alpha$ -pinene and  $\rho$ -cymene being dependent on light and temperature, analogous to isoprene. 678 However, during the night both monoterpenes were also present, albeit at lower mixing ratios, and the 679 nocturnal mixing ratios of the monoterpenes did not decrease as much as isoprene. This has also been 680 noted in previous studies (Yáñez-Serrano et al., 2015).

681

682Despite the higher mixing ratios of limonene compared to other monoterpene species (other than683α-pinene), it was not possible to distinguish any clear diel pattern in the average data for this species684(see Figure 2).  $\beta$ -Ppinene and α-terpinene likewise showed no obvious diel pattern in the rainforest air,685but were found to be greater thanabove the detection limit of the GC-FID of 2 ppt.

686

687 <u>On the contraryIn contrast to plant species of cooler climates, such as spruce, which emit ter-</u>
 688 penes from pools (Ghirardo et al., 2010; Lerdau et al., 1997), Amazonian plant species have been found
 689 to show an emission dependency on light and temperature (Bracho-Nunez et al., 2013; Jardine et al.,

690	2015; Kuhn et al., 2002, 2004). This could partly explain the diurnal pattern of $\alpha$ -Ppinene mixing ratios,
691	which exhibit some relation to a light and temperature dependent emission flux (Kuhn et al., 2002;
692	Rinne et al., 2002; Williams et al., 2007). However, this behaviour wasis not observed for all monoter-
693	pene species. Therefore, the observed diurnal cycles of some monoterpene species might be triggered
694	byrelated to a stronger temperature affects response.
l 695	
696	3.2. Chemodiversity
697	The chemical speciation (or chemodiversity) of monoterpenes relates to the relative abundances
698	of the different monoterpene species in the sampled air. $\alpha$ -Pinene, limonene, myrcene, $\rho$ -cymene and $\beta$ -
699	pinene represented more than 85% of the total monoterpeneMT mixing ratio (Figure 3). During the day
700	(0900h to 1700h) $\alpha$ -pinene had an average abundance (average-±standard deviation) of 46±25% and
701	36±4% of the total monoterpene mixing ratios at 24 and 12 m, respectively, and was the dominant mon-
702	oterpene in this study overall. However, during the night (2000h to 0500h), its relative abundance
703	dropped to $25\pm13\%$ and $25\pm9\%$ at 24 and 12 m, respectively. In contrast, limonene made up $23\pm15\%$
704	and 20±10% of the monoterpenes at 24 and 12 m, respectively, by day, and increased during night time
705	to 33±15% and 26±16% at 24 and 12 m. Thus, there wasis a tendency of towards some differences in
706	monoterpene species abundances between day and night. These wereare mainly due to the nocturnal de-
1 707	creases in $\alpha$ -pinene and the nocturnal relative increase in limonene. It is plausible that the observed de-
708	crease in $\alpha$ -pinene mixing ratios could be due to $\frac{1}{\alpha}$ decreased vegetation emission, as reduced chemical
709	destruction due to very low OH concentrations at night, would lead to an increase in the nocturnal $\alpha$ -
- 710	pinene mixing ratios.

Even though there were clear differences between the absolute and relative abundances of some monoterpene species during day and night, there were no clear changes in the in the vertical gradients (e.g. for  $\alpha$ -pinene night time averages were  $0.15\pm0.05$  ppb for 12 m and  $0.11\pm0.06$  ppb at 24 m). For the day, the apparent difference in the abundance of  $\alpha$ -pinene was due to a single outlier data point covering 30 minutes at noon on 19 October 2015 at 24 m, when the  $\alpha$ -pinene mixing ratio doubled. This increase could not be explained, although it could be related to a strong change in wind speed an hour before the measurement, when the wind was blowing from the North. In general, our observations indicate that the abundance of monoterpene species does not vary much over the heights selected (12 and 24 m) within the canopy. This is consistent with the results by Kesselmeier et al. (2000), where the monoterpene composition at the rain forest floor was comparable to the above-canopy composition at their site.

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### 3.3. Reactivity

724 The variability of the oxidants  $(OH, O_3 \text{ and } NO_3)$  present in the Amazon air is important when 725 considering the impact that monoterpenes can have oin the oxidative regime in the Amazon region and 726 Brazil in general. Hydroxyl radicals are produced mainly during the day via ozone photolysis. and O<sup>4</sup>D 727 reaction with water. Low levels of OH can be also generated by the reaction of ozone with doubly 728 bonded species (e.g. monoterpenes and sesquiterpenes) even at night. In this assessment, we considered 729 the monoterpene contributions to OH reactivity by day only. In contrast, NO<sub>3</sub> is photolytically de-730 stroved during the day, but can become significant at night, so we assessed the impact of monoterpenes on NO<sub>3</sub> reactivity at night. Even though in the Amazon rainforest ozone levels are low (~10-20 ppb) 731 732 compared to other areas of the world (e.g., Williams et al., 2016), it is nevertheless present, and some monoterpenes are extremely reactive towards ozone. Table 2 gives an overview of the lifetime and reac-733 734 tivity (which is defined as reaction rate constant (oxidant i.e. OH)\*[monoterpene species]) to 1 ppb of all the investigated monoterpene species for these three oxidants. For calculating the lifetime of the dif-735 736 ferent monoterpenes as presented in Table 2, typical oxidant concentrations for the Amazon rainforest conditions were used. For OH a mean value of  $7x10^5$  molecules cm<sup>-3</sup> was used as representative of the 737 738 site (Spivakovsky et al., 2000). For ozone reactivity calculations, 12 ppb was used, as this mixing ratio was observed during the measurement period. NO<sub>3</sub> mixing ratios were taken from the MLC-CHEM 739 740 model simulations that predicted mixing ratios of  $\sim 0.4$  ppt.

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Despite, While the most abundant species were α-pinene, limonene and myrcene were the most
 abundant species, - However, with respect to their reactivities towards the different oxidants, their relative contribution to total monoterpene reactivity dramatically changed was not proportional to their

745abundances. The most abundant monoterpene,  $\alpha$ -pinene was not the dominantting sink for the oxidants.746For instanceIn particular,  $\alpha$ -terpinene dominated ozone reactivity associated with monoterpene abun-747dance both during the day and night, as well as the nocturnal nitrate reactivity, despite the low mixing748ratios measured for this compound (Table 2).

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The monoterpene ozone reactivity wasis comparable between day  $(1.37 \times 10^{-6} \text{ s}^{-1})$  and night  $(1.12 \times 10^{-6} \text{ s}^{-1})$ 750 <sup>1</sup>). α-Tterpinene dominated the monoterpene-ozone chemistry, followed by myrcene and limonene. De-751 752 spite the relatively high abundance of  $\alpha$ -pinene (46±25%; average mixing ratio and standard deviation during the day was  $0.34\pm0.04$  ppb at 12 m), its contribution to ozone reactivity with respect to other 753 754 monoterpene species was only  $11\pm7$  % and  $3\pm1$  % at 24m, during the day and night, respectively, at 24 755  $m_{\tau}$  and  $2\pm 1$  % for both day and night at 12 m (Figure 4). As previously noted, the differences in ozone 756 reactivity between heights were negligible for the night and slightly higher at 24 m during the day. As ozone mixing ratios are quite similar for both heights during day and night (11.4 ppb at 12 m and 10.4 757 ppb at 24 m during night, and 16.1 ppb at 12 m and 15.6 at 24 m during the day), the higher abundance 758 759 of  $\alpha$ -pinene during the day, and the lower  $\alpha$ -terpinene mixing ratios at 24 m during the day mainly ex-760 plain these changes in monoterpene-ozone reactivity. It is important to note that these results are de-761 rived from a relative abundance analysis, and unmeasured monoterpene species could change the pro-762 portions, although given the close similitude between PTR-MS and GC-FID measurements shown in Figure 1 this is unlikely. On the other hand, very reactive species which could dominate reactivity, may 763 764 be present in very low concentrations, and which our measurements capabilities would not allow for its 765 monitoringbeing detected.

| 766

The monoterpene reactivity towards <u>the NO<sub>3</sub></u> radical during the night <u>in this study was</u> also dominated by  $\alpha$ -terpinene (<u>40±36%</u> and <u>42±27</u>%, <u>respectively</u> for 24 and 12 m, <u>respectively</u>), although contributions of limonene (30±13% and <u>25±14</u>%, <u>respectively</u> for 24 and 12 m, <u>respectively</u>),  $\alpha$ -pinene (<u>11±6</u> and <u>11±4</u>%, <u>respectively</u> for 24 and 12 m, <u>respectively</u>), and myrcene (<u>13±11</u> and 1<u>6±1</u>2%, <u>re-</u> <del>spectively</del> for 24 and 12 m, respectively</u>) were also significant. No significant differences between the reactivit<u>es</u> at different heights <u>were observed</u>, suggesting a rather homogeneous chemical regime regarding monoterpene chemical destruction within the canopy (from 12 to 24 m). However, note that this finding reflects the use of a single simulated NO<sub>3</sub> <u>mixing ratio</u> due to the absence of direct measurements in the Amazon rainforest, which prevents us from drawing any further conclusion. <u>The Our OH</u> reactivity estimates demonstrate<u>d</u> the important role of myrcene with its higher reactivity towards OH due to its acyclic nature, especially at 12 m<sub>a</sub> where myrcene <u>was</u> more abundant. The total OH reactivity for the sum of monoterpenes was calculated to be 2.<u>4</u> and 3.4 s<sup>-1</sup> for 24 and 12 m, respectively.

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780 As demonstrated in this data set, chemically speciated measurements are very important for understanding how monoterpenes affect Amazon air chemistry dependent on time of day and season, as 781 each monoterpene species has a different reactivity. Therefore, a lower abundance of a certain monoter-782 pene species could not necessarily be related to a lower vegetation emission, but also to a higher 783 reactivity with atmospheric oxidants. Despite the small amount of  $\alpha$ -terpinene present in the atmos-784 phere, it can profoundly affect reactivity due to its fast reaction rate (its lifetime, according to the oxi-785 dant mixing ratios stated above, can be 103, 2 and 11 minutes to OH, O<sub>3</sub> and NO<sub>3</sub>, respectively, -{Neeb 786 et al., 1997). In terms of total OH reactivity accounted for by the monoterpenes, the values of this study 787 788 are very low compared to the total OH reactivity measurements by Nölscher et al. (2016), with a mean of total OH reactivity for the dry season of 32 s<sup>-1</sup>, mostly dominated by isoprene chemistry. This sug-789 790 gests that the monoterpenes contributed only a small fraction to the total OH reactivity at the ATTO site during the investigated time period. This study demonstrates that the abundance does not relate to the 791 importance in chemical reactivity, and species that are usually not considered by atmospheric chemistry 792 793 models due to their modest mixing ratios might actually play a dominant role in the monoterpene atmospheric chemistry. Therefore, it is questionable to generalize the representation of terpene chemistry 794 795 in models (Hallquist et al., 2009) using one or two monoterpene species only. 796

The gas-phase oxidation of the monoterpenes in the Amazon has numerous impacts on the envi ronment, including the production of a multitude of new compounds that are generally longer lived than

799 the primary emissions, increasing the lifetimes and particle production potential of certain compounds by suppressing oxidant availability. Moreover, production of OH due to the ozonolysis of monoterpenes 800 is known to occur (Paulson et al., 1999). The production strength varies depending on the position of 801 the double bonds, if there is more than one (Herrmann et al., 2010). Furthermore, the products of the 802 reaction can be manifold. For instance, when  $\alpha$ -pinene is oxidized by OH, especially at low nitrogen ox-803 804 ides mixing ratios, pinonaldehvde is formed in high vields (Eddingsaas et al., 2012). Chemical processing of  $\alpha$ -pinene can also result in a further production of different monoterpenes such as the reaction 805 of  $\alpha$ -pinene with nitrate during the night, which can lead to the formation of  $\alpha$ -cymene (Gratien et al., 806 807 2011).

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809 The implications of the measured monoterpene abundances for SOA formation at the ATTO site are difficult to quantify, . This is caused by because the SOA formation yield being is dependent on many 810 factors. For example, it depends on the pre-existing organic aerosol mass- into which these products can 811 812 be absorbed (Griffin et al., 1999), and thus the SOA yield can vary. This is the reason why in between regions with similar monoterpene mixing ratios and different aerosol mass loadings, the SOA yield can 813 vary. It also varies strongly between different oxidants and terpene species. For instance,  $\alpha$ -pinene 814 815 forms negligible aerosol mass under NO<sub>3</sub> oxidation (Fry et al., 2014), whereas there is production of organic aerosols when the oxidation of  $\alpha$ -pinene involves O<sub>3</sub> (Ehn et al., 2014) and OH (Eddingsaas et al., 816 817 2012). Monoterpenes containing endocyclic double bonds (e.g.  $\alpha$ -pinene, 3-Carene) or open chains (e.g. 818 myrcene) tend to form less aerosols mass from ozonolysis than monoterpenes with exocyclic double 819 bonds (e.g. β-pinene, sabinene, Hatakeyama et al., 1989; Hoffmann et al., 1997). Following the equation 820 established by Bonn et al. (2014, Eq. 5 in text), we were able to estimate the potential aerosol particle number formation rate initiated by monoterpene species only  $(1 \times 10^{-5} \text{ to } 5 \times 10^{-5} \text{ cm}^{-3} \text{ s}^{-1} \text{ at } 24 \text{ m})$  assum-821 ing steady state conditions for radicals. Those were found to be approximately two orders of magnitudes 822 smaller than the calculated potential new aerosol particle formation rate caused by oxidation products of 823 824 sesquiterpenes. Our calculations assume mixing ratios of sesquiterpenes of 0.2 ppb revealing potential formation rates of (1x10<sup>-3</sup> and 4.5x10<sup>-3</sup> cm<sup>-3</sup> s<sup>-1</sup> at 24 m) based on previous measurements in the Ama-825 zon (Jardine et al., 2011), which are remarkably smaller than observed at mid-latitude conditions (Bonn 826

827 et al., 2014). Furthermore, the level of <del>present</del>-NO present (nitric oxide) also affects -severely the poten-828 tial aerosol growth (Wildt et al., 2014) and yield (Sarrafzadeh et al., 2016) at low BVOC/NO<sub>x</sub> ratios). 829 As the theory assumes contributions of larger organic peroxy radicals ( $RO_2$ ), which are destroyed by 830 reactions, e.g. with NO, increasing NO<sub>x</sub> at constant BVOC mixing ratio will decrease the BVOC/NO<sub>x</sub> ratio and lead to a decline in SOA yield. Our calculations showed this effect, with a change of NO from 831 832 0.2 ppbvppb to 1 ppbvppb leading to a decrease in the formation rate at a diameter of 3 nm. This interdependence calls for a consistent consideration of the BVOC and NO<sub>x</sub> exchange in aerosol formation 833 834 and growth studies.

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#### 3.4. Seasonality

837 By examining GC-FID data collected in previous campaigns, an intra- and inter-annual compari-838 son can be made. These earlier samples were collected using a GSA SG10-2 personal pump sampler. Adsorbent tubes were filled at 167 cm<sup>3</sup>-min<sup>-1</sup> (STP) air flow for 20 min. Total monoterpene averages for 839 840 each season were calculated from 1100h to 1600h LT-at 24 m. Based on these data, we distinguished the monoterpene mixing ratios representative for the dry season, the wet season and the wet-to-dry tran-841 842 sition. The dry season conditions were represented by measurements collected in November 2012, Sep-843 tember 2013, August 2014, and the measurements from this study in October 2015. The wet season 844 measurements were collected in March 2013 and the wet-to-dry transition measurements were collected 845 in June 2013. For the dry season conditions, the total monoterpene mixing ratios were substantially 846 higher (1.02 ppb) compared to the observed monoterpene mixing ratios in the wet season (0.14 ppb) and 847 the wet-to-dry transition season (0.18 ppb) (Figure 5). This coincides with the occurrence of the highest 848 radiation levels and temperatures as well as the lowest precipitation during these dry season measure-849 ment campaigns. During the wet season, the total monoterpene mixing ratios wereare lowest, while dur-850 ing the transition season in June, they wereare slightly higher.

For each season, an average monoterpene chemodiversity distribution is shown in Fig. 5. During the dry seasons, the chemodiversity seems relatively similar ( $39.4\pm4\%$  for  $\alpha$ -pinene,  $20.3\pm3\%$  for limonene), whereas it slightly changes during the wet season, and dramatically changes during the wet-todry transition. The reason for this difference in June could be related to changes in the phenology, as demonstrated at a Central Amazonian site (Alves et al., 2016; Lopes et al., 2016). Furthermore, during the dry season of 2015 a very strong El\_-Niño event <u>was takingtook</u> place, leading to extremely dry conditions observed region-wide (Jardine et al., 2017).

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It has been shown previously that the amounts and speciation of monoterpenes vary strongly ac-860 861 cording to plant species and leaf developmental stage. For instance, Bracho-Nunez et al. (2011) found voung leaves of the some Mediterranean plant species to emit more  $\alpha$ -pinene and mature leaves to emit 862 *e*-ocimene, *z*-ocimene and myrcene, but not *a*-pinene. Some species have been found to be higher emit-863 ters of  $\alpha$ -pinene (i.e. *Hevea spruceana*), whereas others are higher emitters of myrcene (i.e. *Quercus*) 864 *coccifera*, Bracho-Nunez et al., 2013). The leaf developmental stage is also important, as reported for 865 flushing young leaves emitting monoterpenes, in contrast to the isoprene emission of mature leaves of 866 the same plant species (Kuhn et al., 2004). Such a behaviour could explain the lower mixing ratios and 867 different chemodiversity found in June. During this time of the year, leaf flushing takes place in the 868 Central Amazon region (Alves et al., 2016; Lopes et al., 2016). Under these conditions, lower  $\alpha$ -pinene 869 mixing ratios were found as compared to the dry season, when young leaves reach mature levels. There-870 fore, the seasonality in Amazon forest monoterpene emissions might depend more on the changes in ag-871 gregated canopy phenology than on the seasonality of climate drivers (Wu et al., 2016). Our study 872 shows that chemodiversity remains relatively constant during at least the dry seasons, but changed be-873 tween different seasons. Therefore, the implications to the atmosphere are different for each monoter-874 pene species. Kesselmeier et al. (2002) also showed this type of behaviour in their study, where they did 875 876 not find a strong difference in total mixing ratios, but different chemodiversity between seasons, likely expressing differences in seasonal plant developments and atmospheric reactivities, which should be 877 878 accounted for in model implementations at the ATTO site. 879

# 3.5. Modelling analysis

To further support our analysis of the observed magnitude as well as temporal variability in the monoterpene mixing ratios inside the forest canopy, we used MLC-CHEM to: 1) explore how well the model represents the measured mixing ratios and 2) to assess the role of the different in-canopy processes in explaining the diel cycle of the observed monoterpene mixing ratios at the ATTO site.

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886 From Fig. 6, which shows a comparison of the simulated (12.5 and 22.5 m) and observed (12 and 24 m) speciated monoterpene mixing ratios from 17 to 20 October 2015, it can be inferred that the 887 simulated speciated monoterpene mixing ratios are of comparable magnitude to the measured observa-888 tions. This comparison regarding the magnitude of observed and simulated mixing ratios isserves 889 mainly of relevance to appreciate assess the validity of the required selection of basal emission fluxes for 890 891 the different monoterpene compounds. A more relevant result seems to be the overall quite good agree-892 ment between the simulated and observed temporal variability in monoterpene mixing ratios. Note that 893 we also conducted a simulation in which we applied temperature and light dependent monoterpene emission flux. However, those simulations did not follow the observed magnitudes and temporal varia-894 bility as well as the model simulations considering monoterpene emissions that only depend on temper-895 896 ature.

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The generally quite good agreement between the simulated and observed monoterpene mixing 898 ratios, except of an overestimation of simulated  $\alpha$ -pinene mixing ratios for the 17<sup>th</sup> of October. ex-899 presses the overall result of temporally varying emissions, in-canopy chemistry, turbulent mixing and 900 deposition. The latter also involves a potentially important role of deposition to wet leaf surfaces (the 901 inferred wet surface uptake resistances for the monoterpenes are ~300 s m<sup>-1</sup>, similar to values reported 902 by Zhou et al., (2017), MLC-CHEM uses relative humidity as a proxy for the fraction of the leaf surface 903 being wet, (Lammel, 1999; Sun et al., 2016). This results in substantially smaller estimates of canopy 904 905 wetness on 17 October 2015-compared to the following days, which partly explains the simulated high

906	<u><math>\alpha</math>-pinene mixing ratios</u> . The simulated $\alpha$ -pinene mixing ratios for 18-20 <sup>th</sup> -of October, with inferred wet
907	surface fractions up to 1 during the night and ~0.5 during daytime, are in much better agreement with
908	the observations. Regarding the comparison of the simulated observed mixing ratios for some of the
909	other monoterpenes, the simulated ß-pinene, limonene, and myrcene mixing ratios, especially at 12.5 m
910	seem to capture the observed temporal variability quite well. Note that this result for limonene reflects
911	the use of a high leaf basal emission factor (0.9 µg C g <sup>-1</sup> hr <sup>-1</sup> ) required to simulate mixing ratios reach-
912	ing up to 0.4 ppb. These MLC-CHEM simulations were also used to infer how much of the actual emis-
913	sion flux escapes the canopy, expressed by the calculated atmosphere-biosphere limonene flux divided
914	by the canopy emission flux of limonene. This ratio reaches a maximum value of 0.5 around noontime,
915	implying that these model simulations indicate that at the middle of the day, about 50% of the emitted
916	limonene is removed inside the canopy by in-canopy oxidation -and deposition. During night time, this
917	ratio reaches a minimum < 0.1 indicating simulation of very efficient in-canopy removal.

919 These modelling results should be interpreted with carecaution, also given that some of the sim-920 ulated processes cannot be evaluated due to missing observations of canopy wetness as well as the uptake efficiency of monoterpenes by wet surfaces. It should be considered that the simulated removal of 921 922 monoterpenes by wet canopy surfaces could also compensate for a misrepresentation of other canopy 923 processes, e.g., reduced emissions from wet canopy surfaces or an underestimation of the oxidation effi-924 ciency. Further analysis of the model simulated process tendencies (Ganzeveld et al., 2008) indicates 925 only small changes in the simulated source of the monoterpenes over the 4-day period. Regarding the 926 sink of, for example,  $\alpha$ -pinene, chemical destruction of  $\alpha$ -pinene oxidation by O<sub>3</sub>, OH and NO<sub>3</sub> appears to be a relative small term, with the overall sink being dominated by deposition to wet surfaces showing 927 quite large temporal variability. Consequently, the presented quite reasonable agreement between simu-928 lated and observed temporal variability in monoterpenes mixing ratios indicates that deposition to wet 929 930 surfaces may play an important role in monoterpene atmosphere-biosphere exchange. -This should be 931 further corroborated, calling for experiments to determine the actual efficiency (and mechanisms) of up-932 take of monoterpenes by wet canopy surfaces.

### 933 4. Conclusions

934	This study presents an analysis of the measured monoterpene chemodiversity at the Amazon
935	tropical forest measurement site, ATTO. The results showed a distinctly different chemical speciation
936	between day and night, whereas there were little vertical differences in speciation within the canopy (12
937	and 24 m). Furthermore, inter- and intra-annual results demonstrate similar chemodiversity during the
938	dry seasons analysed, but this a change of chemodiversity changed with season, similar to the seasonal
939	measurements performed by Kesselmeier et al. (2002). Furthermore, reactivity calculations demon-
940	strated that higher abundance of a monoterpene species MT does not automatically imply higher reac-
941	tivity, as the most abundant compounds may not be the most atmospheric chemically relevant com-
942	pounds, or the relative contribution of different monoterpenes may change with time. Our calculations
943	support the view to-that the role of canopy exchange may be erroneously estimated when not taking into
944	account speciation-based reactivity in models. Moreover, simulations with a canopy exchange model-
945	ling system to assess the role of canopy interactions compared relatively well with the observed tem-
946	poral variability in speciated monoterpenes, but also indicate the necessity of more experiments to en-
947	hance our understanding of in-canopy sinks of these monoterpenescompounds.

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# 949 5. Data Availability

Even though the data are still not available in any public repository, the data are available upon request from the main author.

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969 authors and not of the participating institutions.

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- **1158** Table 1: <u>Average Mmixing ratio with standard deviation s</u>-in ppb at 24 and 12 m of the measured monoterpene species
- 1159 fromer 17 to 20 October 2015 as determined by the GC-FID analysis. The daytime period was considered chosen from
- 1160 0900h to 1700h and the night time period from 2000h to 0500h (Local time). BLD stands for below detection limit.

Compound	Day 12 m	Night 12 m	Day 24 m	Night 24 m	
α-Pinene	$0.33\pm0.04$	0.15±0.05	0.38±0.21	0.11±0.06	
Limonene	0.18±0.09	$0.18 \pm 0.10$	0.19±0.12	$0.14 \pm 0.07$	
Myrcene	0.16±0.14	$0.12 \pm 0.09$	$0.09 \pm 0.04$	$0.07 \pm 0.06$	
P-Cymene	$0.07 \pm 0.03$	$0.04 \pm 0.01$	$0.08 \pm 0.04$	$0.04 \pm 0.02$	
β-Pinene	0.08±0.03	$0.06 \pm 0.03$	$0.05 \pm 0.03$	$0.04 \pm 0.02$	
Camphene	0.03±0.03	$0.02 \pm 0.01$	$0.03 \pm 0.02$	$0.01 \pm 0.01$	
α-Terpinene	$0.03 \pm 0.02$	$0.03 \pm 0.02$	$0.01 \pm 0.02$	$0.02 \pm 0.02$	
γ-Terpinene	$0.02 \pm 0.01$	$0.01 \pm 0.01$	$0.01 \pm 0.01$	$0.01 \pm 0.01$	
3-Carene	0.001±0.003	$0.003 \pm 0.008$	$0.003 \pm 0.011$	0 or BLD	
α-Phellandrene	0 or BLD	0 or BLD	0 or BLD	0 or BLD	
Sabinene	0 or BLD	0 or BLD	0 or BLD	0 or BLD	
MT Sum – GC-FID	0.91±0.10	0.62±0.19	0.82±0.34	0.45±0.13	
MT Sum – PTR-MS	0.96±0.27	$0.54 \pm 0.17$	$0.77 \pm 0.22$	$0.56 \pm 0.16$	

**1162** Table 2: Lifetime of the different monoterpene species related to OH, O<sub>3</sub> and NO<sub>3</sub> for the OH daytime conditions at 24 m and at 12 m. In addition, the normalized reactivity to 1 ppb of the different monoterpene species is calculated.

Monoterpenes investi-	Formula	Lifetime (minutes)			<u>Normalized r</u> Reactivity to 1 ppb s <sup>-1</sup>		
gated		OH	<b>O</b> <sub>3</sub>	$NO_3$	OH	<b>O</b> <sub>3</sub>	NO <sub>3</sub>
α-Pinene	$C_{10}H_{16}$	449	615	250	1.42	2.3E-06	0.17
Camphene	$C_{10}H_{16}$	447	57422	2461	1.43	2.4E-08	0.02
Sabinene	$C_{10}H_{16}$	400	623	155	1.60	2.2E-06	0.27
β-Pinene	$C_{10}H_{16}$	320	3445	618	2.00	4.0E-07	0.07
Myrcene	$C_{10}H_{16}$	71	110	141	8.98	1.3E-05	0.30
α-Phellandrene	$C_{10}H_{16}$	132	17	21	4.84	8.1E-05	1.96

∆3-Carene	$C_{10}H_{16}$	271	1397	170	2.37	9.9E-07	0.24
α-Terpinene	$C_{10}H_{16}$	103	2	11	6.24	5.6E-04	3.76
ρ-Cymene	$C_{10}H_{14}$	1577	>90000	>90000	0.41	1.3E-09	2.7E-05
Limonene	$C_{10}H_{16}$	145	246	127	4.41	5.6E-06	0.33
γ-Terpinene	$C_{10}H_{16}$	140	369	53	4.57	3.8E-06	0.78
Isoprene	C <sub>5</sub> H <sub>8</sub>	238	4069	238	2.69	3.4E-07	0.02





Figure 1: Graph showing the speciated monoterpene mixing ratios measured hourly from 17 to 20 October 2015 for-b) 24 m(b) and-c) 12 m(c). The colours on the stacked bar plot indicate the different monoterpene species as they are denoted in the legend. The black line represents the PTR-MS total monoterpene mixing ratio, with a gap of data on the 19 October 2015. Temperature at 80 m is shown in-as the red thick line and photosynthetically active radiation at 39 m is shown in-by the shaded areas in-(a).



Figure 2: Average diel cycles for α-pinene (a), limonene (b), myrcene (c), ρ-cymene (d), β-pinene (e) and α-terpinene (f) mixing ratios for 24 m (dashed line) and 12 m (thick line). In the back, average diel cycle of isoprene mixing ratios as measured by the GC-FID are shown for 24 m (light green) and 24 m (dark green). Error bars represent the standard deviation of the averages.

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1179 Figure 3: Pie charts representing the day (a and c) and night (b and d) averaget monoterpene species

- abundances from 17 to 20 October 2015, inwith the average percentages and standard deviations at 24
- 1181 (a and b) and 24 (c and d) m. DThe day period was from 0900h to 1700h and the night period was from
- 1182 <u>2000h to 0500h.</u>
- 1183



Figure 4: Pie charts representing day (a and e) and night (b and f) ozone reactivity, OH reactivity (only for day, c and h) and NO<sub>3</sub> reactivities from 17 to 20 October 2015, with the average percentages and

standard deviations (only for night, d and i), for 12 m on the bottom and 24 m on the top. The day period was from 0900h to 1700h and the night period was from 2000h to 0500h.





Figure 5: Monoterpene mixing ratio chemical speciation during the seasons of measurements. In the top, the monthly average of temperature (in red) and photosynthetically active radiation (in orange) are displayed with their standard deviations for the 80 m height. Rain, also on top, is displayed as mm per month (bars). In-At the bottom, the different monoterpene species are differentiated by colours, stacked together adding up to the sum of monoterpenes. On top of each bar, a chart pie with the chemical speciation is shown for easier visualization.

