1	Using different assumptions of aerosol mixing state and chemical
2	composition to predict CCN concentrations based on field
3	measurements in urban Beijing
4 5	Jingye Ren ¹ , Fang Zhang ^{1, 2*} , Yuying Wang ¹ , Don Collins ³ , Xinxin Fan ¹ , Xiaoai Jin ¹ , Weiqi Xu ^{3, 4} , Yele Sun ^{3, 4} , Maureen Cribb ⁵ , Zhanqing Li ^{1, 5}
6	
7 8	¹ College of Global Change and Earth System Science, Beijing Normal University, Beijing 100875, China
9	² Joint Center for Global Change Studies (JCGCS), Beijing 100875, China
10 11	³ Department of Atmospheric Sciences, Texas A&M University, College Station, TX, USA
12	⁴ State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric
13	Chemistry, Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing
14	100029, China
15	⁵ University of Chinese Academy of Sciences, Beijing 100049, China
16	⁶ Earth System Science Interdisciplinary Center and Department of Atmospheric and
17	Oceanic Science, University of Maryland, College Park, Maryland, USA
18	
19	
20	
21	
22	*Correspondence to: Fang Zhang (fang.zhang@bnu.edu.cn)
23	
24	1

25 Abstract

26	Understanding the impacts of aerosol chemical composition and mixing state on
27	cloud condensation nuclei (CCN) activity in polluted areas is crucial for accurately
28	predicting the CCN number concentrations (N _{CCN}). In this study, we predict N _{CCN}
29	under five assumed schemes of aerosol chemical composition and mixing state based
30	on field measurements in Beijing during the winter of 2016. Our results show that the
31	best closure is achieved with an assumption of a size dependent chemical composition
32	for which sulfate, nitrate, secondary organic aerosols and aged black carbon are
33	internally mixed with each other but externally mixed with primary organic aerosol
34	and fresh black carbon (external-internal size-resolved, abbreviated as EI-SR scheme).
35	The resulting ratios of predicted-to-measured N _{CCN} ($R_{CCN_p/m}$) were 0.90–0.98 under
36	both clean and polluted conditions. Assumption of an internal mixture and bulk
37	chemical composition (INT-BK scheme) shows good closure with $R_{CCN_p/m}$ of
38	1.01–1.16 under clean conditions, implying that it is adequate for CCN prediction in
39	continental clean regions. On polluted days, assuming the aerosol is internally mixed
40	and has a chemical composition that is size dependent (INT-SR scheme) achieves
41	better closure than the INT-BK scheme due to the heterogeneity and variations in
42	particle composition at different sizes. The improved closure achieved using the
43	EI-SR and INT-SR assumptions highlight the importance of measuring size-resolved
44	chemical composition for CCN predictions in polluted regions. N_{CCN} is significantly
45	underestimated (with $R_{CCN_p/m}$ of 0.66–0.75) when using the schemes of external
46	mixtures with bulk (EXT-BK scheme) or size-resolved composition (EXT-SR

47	scheme), implying that primary particles experience rapid aging and physical mixing
48	processes in urban Beijing. However, our results show that the aerosol mixing state
49	plays a minor role in CCN prediction when the κ_{org} exceeds 0.1.
50	1 Introduction
51	Atmospheric aerosol particles can serve as cloud condensation nuclei (CCN) and,
52	in turn, affect the optical and microphysical properties of clouds (Twomey, 1977;
53	Albrecht, 1989; Charlson et al., 1992). Additionally, an increase in the aerosol number
54	concentration may suppress precipitation in shallow clouds and promote it in deep
55	convective clouds (Rosenfeld et al., 2008; Li et al., 2011). A key challenge to
56	understanding indirect aerosol effects is quantifying CCN spectra and their spatial and
57	temporal variations.
58	The ability of particles to act as CCN mainly depends on their size, chemical
59	composition, and mixing state (McFiggans et al., 2006; Dusek et al., 2006; Ma et al.,
60	2013). The impacts of the size distribution and chemical composition on CCN activity
61	has been discussed in previous studies (Dusek et al., 2006, Ervens et al., 2007;
62	Broekhuizen et al., 2006; Yum et al., 2005, 2007; Wiedensohler et al., 2009; Deng et
63	al., 2013; Zhang et al., 2014, 2016; Kawana et al., 2016). The effect of chemical
64	composition can be represented by a hygroscopicity parameter (κ) (Petters and
65	Kreidenweis, 2007) that is often used to predict N _{CCN} (Moore et al., 2012; Zhang et al.,
66	2014). However, particle composition may vary from single species to a mixture of
67	multiple species for a given size. A description of size-resolved chemical composition

68	thus leads to a better prediction of N_{CCN} because it allows variation of κ with size
69	(Medina et al., 2007; Wang et al., 2010; Meng et al., 2014). Variations in mixing state
70	also impact $N_{\mbox{\scriptsize CCN}}$ prediction, with the effect dependent on the hygroscopicity of the
71	organic component (Wang et al., 2010). The assumption of internal mixtures has been
72	demonstrated to predict N_{CCN} well (Ervens et al., 2007; Chang et al., 2007; Andreae
73	and Rosenfeld, 2008; Gunthe et al., 2009; Rose et al., 2008; Meng et al., 2014; Zhang
74	et al., 2014; Li et al., 2017). However, some studies have shown that detailed
75	information about the chemical composition and the mixing state is required because
76	of the complexity of the hygroscopicity of organics (Broekhuizen et al., 2006; Bhattu
77	and Tripathi, 2015) and the differences in the CCN activity between fresh and aged
78	aerosols (Gunthe et al., 2011). Therefore, the impact of different assumptions
79	concerning the mixing state and chemical composition on accurately quantifying CCN
80	concentrations needs further investigation, especially in heavily polluted regions.
81	Beijing, a typical polluted city, frequently experiences severe haze pollution
82	episodes (Sun et al., 2013; Guo et al., 2014; Zheng et al., 2015), particularly in winter.
83	Several recent studies have focused on studying particle hygroscopicity (Wu et al.,
84	2016; Wang et al., 2017) and chemical composition (Gunthe et al., 2011), and using
85	bulk κ to predict CCN in Beijing (e.g., Liu et al., 2014; Zhang et al., 2017). However,
86	to our knowledge, a comprehensive CCN closure test considering chemical
87	composition and mixing state is lacking for this polluted urban area. In particular, the
88	transformation of the particle mixing state may be very quick during severe pollution
89	conditions (Wu et al., 2016). During pollution events, the hygroscopicity of organics

90	and the CCN activity are often enhanced rapidly with the aging process (Gunthe et al.,
91	2011; Kawana et al., 2016). Therefore, the characterization and parameterization of
92	CCN activation may be more challenging in polluted regions due to the impacts of
93	organics (Wang et al., 2010; Meng et al., 2014; Che et al., 2016; Zhang et al., 2016).
94	In this study, we use size-resolved measurements of CCN activity and
74	In this study, we use size-resolved measurements of every activity and
95	size-resolved chemical composition information to predict N_{CCN} using field
96	measurement data collected in Beijing during the winter of 2016. The CCN closure
97	study is carried out using five schemes with different assumptions of particle mixing
98	state and chemical composition. By classifying the data into three different periods
99	(nighttime, noontime, and the evening rush hour), we also investigate the variations in
100	aerosol mixing state from fresh to relatively aged aerosols. The sensitivity of
101	predicted N_{CCN} to the particle mixing state and organic volume fraction with the aging
102	of organic particles is also presented in the last section of the study.

2 Measurements and data

Data used here were measured from 15 November to 14 December 2016
during the Air Pollution and Human Health (APHH) field campaign at the Institute of
Atmospheric Physics (IAP), Chinese Academy of Sciences (39.97°N, 116.37°E),
which is a typical urban site with influences from traffic and cooking emissions (Sun
et al., 2015). The sampling instruments were placed in a container at ground level.
The particle number size distribution (PNSD) was measured by a Scanning

110	Mobility Particle Sizer (SMPS; Wang et al., 2003). The SMPS consists of a
111	differential mobility analyzer (DMA; model 3081, TSI Inc.) and a condensation
112	particle counter (CPC; model 3772, TSI Inc.). Measurements of size-resolved CCN
113	efficiency spectra were made by an integrated system combining the SMPS (Wang et
114	al., 2003) and a Droplet Measurement Technologies CCN counter (DMT-CCNc;
115	Lance et al., 2006). The procedure to couple the SMPS and the DMT-CCNc
116	developed by Moore et al. (2010) was followed. Atmospheric particles were sampled
117	from an inlet located 1.5 m above the roof of the container and were then passed
118	through a silica gel desiccant drying tube and into the SMPS. The relative humidity
119	of the sample flow was below 30%. The sample flow exiting the DMA was divided
120	into 0.5 lpm for the CCNc and 0.5 lpm for the CPC. Before and after the field
121	campaign ammonium sulfate was used to calibrate the supersaturation (SS) levels of
122	the CCNc with longitudinal temperature differences of 2, 3, 5, 8, 10, 13, and 15 K as
123	shown in Fig. S1. Based on this calibration, the five effective SS levels were 0.12,
124	0.14, 0.23, 0.40, and 0.76%.
125	The PNSD spanned the size range of 10–550 nm with a measurement scan time
126	of 5 min. Total particle or condensation nuclei (CN) size distributions were calculated
127	with the multiple charge correction and transfer function used in the TSI-AIM
128	software. The CN number concentration (N_{CN}) is the total aerosol number
129	concentration and is obtained by integrating the PNSD over the size range of 10–550

nm. The full measurement cycle of the CCNc for the five SS levels took one hour (20

131 min for 0.12% and 10 min for each higher SS). Size-resolved CCN efficiency data

132	were inverted with a multiple charge correction (Moore et al., 2010). The CCN
133	number size distribution was calculated by multiplying the CCN efficiency spectrum
134	by the particle number size distribution. The total CCN concentration was then
135	calculated by integrating the size-resolved N_{CCN} . The bulk activation ratio (AR) was
136	calculated as N_{CCN}/N_{CN} . The results were stratified between polluted and background
137	conditions with an assumed threshold PM1 mass concentration of 50 μ g m ⁻³ .
138	An Aerodyne High-Resolution Time-of-Flight Aerosol Mass Spectrometer
139	(HR-ToF-AMS; DeCarlo et al., 2006) was housed in a sampling room on the rooftop
140	of a two-story building to measure size-resolved non-refractory submicron aerosols,
141	including organics, sulfate, nitrate, ammonium, and chloride with a time resolution of
142	\sim 5 min. More details about the HR-ToF-AMS and the measurement site have been
143	described in previous studies (Sun et al., 2010; Sun et al., 2016). The organics are
144	classified by using Positive Matrix Factorization (PMF) (Paatero and Tapper, 1994),
145	considering as being composed of two components: primary organic aerosol (POA)
146	representing non-hygroscopic particles and secondary organic aerosol (SOA)
147	representing hygroscopic particles. The first component consists manily of
148	hydrocarbon-like organic aerosol (HOA), a surrogate of primary organic aerosol
149	(POA) from local combustion sources. And the size distribution of the primary OA
150	was measured by the estimated size-distribution of the $C_4H_9^+$ fragment (Aiken et al.,
151	2009; Zhang et al., 2005). The size distribution of the SOA was calculated as the
152	difference between the total OA and POA.

153	The black carbon (BC) mass concentration was measured using a
154	seven-wavelength aethalometer (AE33, Magee Scientific Corp.). Zhao et al. (2017)
155	provides details about this instrument and the measurements it makes. Due to an
156	absence of size-resolved BC measurements, the BC size distribution was calculated
157	from the combination of an approximately lognormal distribution measured by a
158	single particle soot photometer (SP2, DMT) (Wu et al., 2017) and the total BC mass
159	concentration. Note that because the SP2 measures BC core diameter instead of the
160	diameter of the BC-containing particle, it would overestimate the BC mass
161	concentration of smaller particles but underestimate that of the larger ones. Such
162	overestimation would likely lead to an underestimation of $N_{\mbox{\scriptsize CCN}}$ due to the increased
163	mass fraction of BC of total particles. The uncertainty of this effect is evaluated in
164	Section 4.3. The fresh and aged BC size distributions are determined from the total
165	BC size distribution measured by the SP2 (Wu et al., 2017) and from the dependence
166	of the fraction of internally mixed soot (F_{in}) on particle diameter (D_p) observed in
167	urban Beijing by Cheng et al. (2012). The instruments produce different diameters. In
168	this paper, we have unified both the aerodynamic diameter from AMS and volume
169	equivalent diameter from SP2 to be mobility diameter. In addition, actual fresh BC
170	particles are not spheres and neither are some of the partially aged BC, but because
171	both the diameter measured from SP2 and the BC size distribution from the literatures
172	are with assumption of the particles being spheres, the fresh and aged BC in this study
173	are thereby assumed to be spherical particles.

3 Theory

175 **3.1 Calculation of CCN concentration using** *k***-Köhler theory**

In this study, we used the critical or cutoff particle diameter (D_{cut}) and particle number size distribution to calculate N_{CCN}. The method to derive D_{cut} is based on κ -Köhler theory (Petters and Kreidenweis, 2007), with the water vapor saturation ratio over the aqueous solution droplet *S* given by:

180
$$S = \frac{D^3 - D_P^3}{D^3 - D_P^3 (1 - \kappa)} \exp(\frac{4\sigma_w M_w}{RT\rho_w D}), \qquad (1)$$

181 where *D* is the droplet diameter, D_p is the dry diameter of the particle, M_w is the 182 molecular weight of water, σ_w is the surface tension of pure water, ρ_w is the density of 183 water, *R* is the gas constant, and *T* is the absolute temperature. When $\kappa > 0.1$ it can be 184 approximately expressed as:

185
$$\kappa = \frac{4A^3}{27D_p^{-3}\ln^2 S_c},$$
 (2)

$$A = \frac{4\sigma_w M_w}{RT\rho_w},\tag{3}$$

187 where S_c is the particle critical supersaturation. The other variables in the equations

are:
$$T = 298.15$$
 K, $R = 8.315$ J K⁻¹ mol⁻¹, $\rho_w = 997.1$ kg m⁻³, $M_w = 0.018015$ kg mol⁻¹,

189 and $\sigma_w = 0.072 \text{ J m}^{-2}$ (Rose et al., 2008).

190 For internally-mixed particles, κ is calculated as follows (Petters and

191 Kreidenweis, 2007; Gunthe et al., 2009):

192
$$\kappa_{chem} = \sum_{i} \varepsilon_{i} \kappa_{i},$$
(4)

193
$$\boldsymbol{\kappa}_{org} = f_{POA} \cdot \boldsymbol{\kappa}_{POA} + f_{SOA} \cdot \boldsymbol{\kappa}_{SOA}, \tag{5}$$

194	where κ_i and ε_i are the hygroscopicity parameter and volume fraction for the
195	individual components in the mixture, and f_{POA} and f_{SOA} are the primary organic
196	aerosol (POA) and secondary organic aerosol (SOA) mass fractions in the mixture.
197	The Aerosol Mass Spectrometer (AMS) mainly measured the particle mass size
198	distributions of SO ₄ ^{2–} , NO ₃ [–] , NH ₄ ⁺ and organic compounds, while the
199	Zdanovskii-Stokes-Robinson relation requires the volume fractions of the particle
200	chemical composition (Stokes and Robinson, 1966; Zdanovskii, 1948). A simplified
201	ion pairing scheme is used to calculate the mass concentrations of the inorganic salts,
202	which includes only NH_4NO_3 and $(NH_4)_2SO_4$ as possible salts (Gysel et al., 2007). In
203	this study, we considered five components: NH4NO3, (NH4)2SO4, SOA, POA, and BC.
204	The $\kappa_{(NH4NO3)}$ is equal to 0.67 and $\kappa_{((NH4)2SO4)}$ is equal to 0.61 (Petters and Kreidenweis,
205	2007; Gunthe et al., 2009). The κ_{org} is estimated using the linear function derived by
206	Mei et al. (2013a), namely, $\kappa_{\text{org}} = 2.10f_{44} - 0.11$, where f_{44} is dependent upon organics
207	oxidation level. The mean κ_{org} is 0.10 in our case. The organics are classified into two
208	factors: POA representing non-hygroscopic particles ($\kappa = 0$) and SOA representing
209	hygroscopic species. In our study, the average contributions of POA and SOA to total
210	organics were 0.53 and 0.47, respectively. On the basis of equation (5), $\kappa_{(SOA)}$ is
211	assumed to be 0.2. Also, $\kappa_{(BC)}$ is assumed to be 0.

212 **3.2** Assumptions about mixing state and chemical composition

213	To examine the influence of the mixing state and chemical composition on
214	CCN activation, five assumptions (Fig. 1) are used to predict N_{CCN} . Although the
215	assumption of completely internal or external mixing for ambient aerosols represents
216	two extremely simplified schemes and may be atmospherically unrealistic, it allows
217	us to understand the importance of the particle mixing state for predicting N_{CCN} . In
218	addition, size independent and dependent compositions are derived from the mass
219	concentrations of different species measured by the AMS so that the impact of
220	chemical composition on CCN activity can be examined. A detailed introduction of
221	the five assumption schemes follows.

Assumption 1: internal mixture with bulk chemical composition (INT-BK)

In this scheme, submicron particles are assumed to be internally mixed with bulk 223 chemical composition, where the mass fraction of each component (e.g. NH₄NO₃, 224 (NH₄)₂SO₄, SOA, POA, and BC) is uniform throughout the full size range as shown in 225 226 Fig. 1a. The overall κ is calculated from the bulk chemical composition measured by the AMS based on the simple mixing rule (Equation 4) to obtain the critical diameter 227 at a given SS. For calculating N_{CCN} all (and only) particles with diameters greater than 228 229 D_{cut} are considered CCN-active. The total N_{CCN} is then calculated from the step-wise integration of the PNSD for $D_p > D_{cut}$. The equations used in the calculations are as 230 follows, 231

$$_{232} \quad CCN_{pre} = \int_{D_{cut}}^{D_{end}} n(\log D_p) d\log D_p \tag{6}$$

$$D_{cut} = \sqrt[3]{\frac{4A^3}{27\sum_i \varepsilon_i \kappa_i \ln^2 S_c}}$$
(7)

where D_{cut} is the critical diameter, D_{end} is the upper size limit of the PNSD, $n (\log D_p)$ is the function of the aerosol number size distribution, *i* is the chemical component element, and the other parameters are the same as those presented in Equations (2), (3) and (4).

238 Assumption 2: internal mixture with size-resolved chemical composition

239 (INT-SR)

For this scheme submicron particles are assumed to be internally mixed and the

241 chemical composition is size-dependent as shown in Fig. 1d. The fractional

contributions of the components at each size bin are derived from mass size

distributions of the five species considered, i.e., NH₄NO₃, (NH₄)₂SO₄, SOA, POA,

and BC.

For this assumption, the critical diameter is derived from the total hygroscopic parameter, κ , at each size bin, *j*. For each size bin for which $D_{p,j}$ is > than the calculated $D_{cut,j}$ the activated fraction was assumed to be 1.0 and for all others it was 0.0. The N_{CCN} is calculated as follows:

$$CCN_{pre} = \int_{D_{begin}}^{D_{end}} n(\log D_p) d\log D_p$$
(8)

$$D_{cut,j} = \sqrt[3]{\frac{4A^3}{27\sum_i \varepsilon_{ij}\kappa_{ij}\ln^2 S_c}}$$
(9)

where D_{begin} and D_{end} are the first and last diameters of the PNSD, n (log D_p) is the function of the aerosol number size distribution, *i* is the chemical component element, *j* is the PNSD size bin, and the other parameters are the same as those presented in Equations (2), (3) and (4).

255 Assumption 3: external mixture with bulk chemical composition (EXT-BK)

For this scheme the submicron aerosol is treated as an external mixture. This means that there are five types of particles, i.e., NH4NO₃, (NH4)₂SO₄, SOA, POA, and BC, and each particle consists of a single component. The volume fraction of each component, which is derived from bulk mass concentrations, does not vary with size (as shown in Fig. 1b).

At a given S, the critical diameter of each particle type is retrieved from the κ of each component. The N_{CCN} of each aerosol type is calculated as the CCN-active particle number concentration multiplied by the bulk volume fraction of the components as expressed in Equation (10). The N_{CCN} of the five particle types are finally summed to obtain the total N_{CCN}. The specific equations are as follows,

266
$$CCN_{pre} = \sum_{i} \left(\int_{D_{icut}}^{D_{end}} n(\log D_p) d \log D_p * V_i \right)$$
 (10)

$$D_{cut,i} = \sqrt[3]{\frac{4A^3}{27\kappa_i \ln^2 S_c}}$$
(11)

where $D_{cut,i}$ is calculated for each component, *i*, at a given SS, V_i is the volume fraction of each aerosol type, n (log D_p) is the function of the aerosol number size distribution, *i* is the chemical component element, and the other parameters are the same as those presented in Equations (2), (3) and (4).

272 Assumption 4: external mixture with size-resolved chemical composition

As with the EXT-BK scheme the same five particle types are considered and 274 their relative concentrations selected to match the measured composition. But unlike 275 276 with the EXT-BK scheme the relative concentrations of the five particle types vary with particle size to capture the size-dependence of the measured composition, as is 277 depicted in Fig. 1e. The volume fraction of each particle type at each size is first 278 multiplied by the total particle number size distribution (PNSD) to get the PNSD_i of 279 each aerosol type. The N_{CCN} of each particle type is then obtained from the step-wise 280 integration of the PNSD_{*i*} for $D_p > D_{cut,i}$, and then summed to get the total N_{CCN}, as 281 described by Equation (12). Similar to EXT-BK, the critical diameter of each particle 282 type is also derived from the κ of each pure component at a given S. 283

$$CCN_{pre} = \sum_{i} \left(\int_{D_{begin}}^{D_{end}} \left(n(\log D_p) * V_{ij} \right) d \log D_p \right)$$
(12)

$$D_{cut,i} = \sqrt[3]{\frac{4A^3}{27\kappa_i \ln^2 S_c}}$$
(13)

where V_i is the volume fraction of each particle type in a size bin, n (log D_p) is the function of the aerosol number size distribution, *i* is the chemical component element, *j* is the particle size bin, and the other parameters are the same as those presented in Equations (2), (3) and (4).

Assumption 5: sulfate, nitrate, SOA and aged BC internally mixed, and POA and fresh BC externally mixed, and all components with size-resolved chemical composition (EI-SR)

293	At each pa	article size	sulfate,	nitrate,	and SOA	with BC-aged	are treated as
-----	------------	--------------	----------	----------	---------	--------------	----------------

internally mixed, but POA and BC-fresh are present in separate particles and are

295 non-hygroscopic. As with INT-SR and EXT-SR the chemical composition is

size-dependent, as shown in Fig. 1c. The EI-SR scheme likely represents a case that is

297 most similar to that of actual atmospheric aerosols in locations such as Beijing. The

298 fresh and aged BC size distributions are determined from the total BC size

distribution measured by the SP2 (Wu et al., 2017) and from the dependence of the

fraction of internally mixed soot (F_{in}) on particle diameter (D_p) observed in urban

301 Beijing by Cheng et al. (2012).

302 In this assumption the fresh BC and POA particles can serve as CCN only if their

diameter is larger than 200 nm; otherwise they are CCN-inactive. Thus, the total N_{CCN} of those externally mixed components (N_{CCN} _EXT) is calculated from the step-wise integration of the product of the PNSD and the volume fraction of the fresh BC and POA in each size bin larger than 200 nm.

The N_{CCN} of the remaining components (sulfate, nitrate, and SOA with BC-aged) that are treated as an internal mixture, denoted as N_{CCN}_INT, is predicted in the same way as for the INT-SR scheme, with the only difference being that the PNSD is first multiplied by the volume fraction of the mixed component particles for each size bin. The total N_{CCN} is thus calculated as the sum of N_{CCN}_EXT and N_{CCN}_INT. The specific equations are as follows,

313
$$CCN_{pre} = \int_{D_{begin}}^{D_{200}} (n(\log D_p) * r_j) d\log D_p + \int_{D_{200}}^{D_{end}} n(\log D_p) d\log D_p$$
 (14)

$$D_{cut,j} = \sqrt[3]{\frac{4A^3}{27\sum_i \varepsilon_{ij} \kappa_{ij} \ln^2 S_c}}$$
(15)

where D_{begin} and D_{end} are the first and last diameters of the PNSD, n (log D_p) is the function of the aerosol number size distribution, *r* is the volume fraction of the internal (hygroscopic) mixture at each size, *i* is the chemical component element, *j* is the particle size bin, and the other parameters are the same as those presented in Equations (2), (3) and (4).

320 4 Results and discussion

4.1 Diurnal variations in aerosol properties

322	Diurnal variations in mean PNSD and bulk chemical composition under
323	polluted and background conditions are shown in Fig. 2. Significant diurnal variations
324	in PNSD are observed during the campaign. For both polluted and background cases
325	the abrupt increases in concentration of small particles ($D_p < 100 \text{ nm}$) from
326	1700–2000 local time (LT) are likely related to fresh primary emissions from cooking
327	and traffic sources (Wang et al., 2017; Zhao et al., 2017), which is also evident in the
328	significant increase in mass concentration of non-hygroscopic POA (Fig. 2d and 2e).
329	The peak amplitude in the PNSD that occurs from about 0800 to1200 LT is probably
330	associated with secondary formation processes, which is indicated by an apparent
331	increase of nitrate, SOA and f_{44} (oxidation level of organics) in the morning (0800 LT)
332	when photochemistry becomes significant. The effect is more apparent on clean days.
333	In addition, the PNSD amplitude and BC and POA concentrations are high at
334	nighttime, suggesting an influence from the diurnal variation of the planetary
335	boundary layer (PBL) height. In particular, on polluted days the PBL plays a key role
336	in regulating the diurnal variation of primary components like POA and BC (e.g.,
337	Dzepina et al., 2009; Cross et al., 2009). On clean days secondary formation and
338	primary sources play dominant roles in regulating diurnal variations. The PNSD in
339	clean cases has peaks at smaller D_p (~30–40 nm, Fig. 1c) compared to polluted cases
340	(~100 nm), which is associated with particle growth accompanying atmospheric
341	chemistry processes during haze evolution (Guo et al., 2014; Wang et al., 2016).

342 **4.2** Cumulative Gaussian distribution function fit and parameters derived from

343 the CCN efficiency

The activated fractions measured at the five supersaturation levels were fitted using the following two functions (Rose et al., 2008; Mei et al., 2013b):

346
$$R_a(S) = \frac{E}{2} \cdot (1 + erf(\frac{\ln S - \ln S^*}{\sqrt{2\sigma_s^2}})), \qquad (16)$$

347
$$f_{N_{CCN}/N_{CN}} = a(1 + erf(\frac{D - D_a}{\sigma_a \sqrt{2}})),$$
 (17)

where Ra(S) and f(NCCN/NCN) are the CCN activation fractions, the maximum activation 348 fraction (MAF) is equal to E or 2a, S^* and D_a are the midpoint activation 349 supersaturation and diameter, respectively, and σ_s and σ_a are the cumulative 350 distribution function (CDF) standard deviations. During this field campaign, 2580 351 352 size-resolved CCN efficiency spectra at five SS levels were measured. To illustrate the characteristics of the activation spectra, the CDF fits are shown in Fig. 3 and in 353 Tables S1-2. A gradual increase in size-resolved AR with SS suggests that particles 354 had different hygroscopicities even at the same diameter. The heterogeneity of particle 355 chemical composition can be represented by the ratio of σ_a and D_a (i.e., σ_a/D_a), where 356 σ_a is the standard deviation derived from the cumulative Gaussian distribution 357 function (Eqn. 12) and D_a is the activation diameter (Rose et al., 2010). The ratio of 358 σ_a/D_a during the three periods is shown in Fig. 3b. 359

360 4.2.1 CCN activation curves and heterogeneity of chemical components

361	For larger particles with $D_p > 100$ nm, no significant differences were observed
362	in the CCN efficiency spectra (Fig. 3a), suggesting a similar hygroscopicity during the
363	three periods. For particles with $D_p < 100$ nm, the CCN efficiency spectrum observed
364	during the evening rush hour period showed a much more gradual increase (with
365	smaller slopes) in size-resolved AR than that derived for the other two periods. This is
366	attributed to the strong influence of POA emissions, which consist of less hygroscopic
367	and externally-mixed smaller particles mainly from cooking and traffic during the
368	evening rush hour period (also indicated by the increased $\sigma_a\!/D_a)$. Particles with $D_p\!<\!$
369	100 nm emitted during the evening rush hour period require a higher SS to reach the
370	same AR. However, for $D_p > 100$ nm the slope of AR with respect to SS was steep
371	and near the instrumental limit obtained for a pure ammonium sulfate aerosol. Che et
372	al. (2016) have reported that particles larger than about 150 nm have relatively
373	uniform composition. This suggests that particles become more internally mixed with
374	growth from the Aitken mode to the accumulation mode. This feature is also
375	suggested by the decreasing σ_a/D_a with increasing particle diameter.

4.2.2 Mean critical activation diameter

The critical activation diameter at different SS levels under background and
polluted conditions is shown in Fig. 4. The difference in critical diameter between
polluted and background cases are calculated as D_p_POL - D_p_BG. At lower SS levels,

380	the critical diameters for polluted cases were slightly smaller than those observed on
381	clean days, suggesting larger particles are more CCN-active on polluted days. This is
382	expected based on HTDMA measurements that showed that particles in the
383	accumulation mode on polluted days are more hygroscopic than those on clean days
384	in urban Beijing (Wang et al., 2017). At higher SS the critical diameter on polluted
385	days was a little higher than that obtained under clean conditions, suggesting that
386	particles with D_p of ~40 nm are less CCN active. This is likely because a high
387	concentration of small and hygroscopic particles in the Aitken mode arise from the
388	photochemistry-driven nucleation process on clean days. However, in polluted cases,
389	small particles are mostly composed of hydrophobic POA from cooking and traffic
390	sources. This was also observed by Wang et al., (2017) who showed that 40 nm
391	particles are less hygroscopic on polluted days. However, the differences in critical
392	diameter between polluted and background cases are small, reflecting a relatively
393	minor influence of hygroscopicity on CCN activity.

394 **4.2.3 MAF**

As shown in Fig. 5, the maximum activated fractions on clean and polluted days during the campaign are less than 1, which suggests that at least some sampled aerosols were externally mixed (Gunthe et al., 2011). For example, the MAF for particles with D_p of ~180 nm was around 0.78 at SS = 0.12% under background conditions, indicating that ~22% of the particles are non-hygroscopic. The higher MAFs under polluted conditions suggest a more internally mixed aerosol (Wu et al.,

401	2016; Wang et al., 2017). The MAF during the 1200–1400 LT period was highest,
402	which is likely due to strong photochemical aging processes that lead to more internal
403	mixing of the aerosol.

404 4.3 CCN closure study and the sensitivity of predicted N_{CCN} to assumed aerosol

405 mixing state and chemical composition

- Fig. 6 shows the comparisons between predicted and measured N_{CCN} at different 406
- SS levels under background and polluted conditions. The ratios of 407
- predicted-to-measured N_{CCN} ($R_{CCN p/m}$) ranged from 0.66 to 1.16, suggesting 408
- significant influences of the different assumptions on CCN prediction. The EI-SR 409

assumption scheme predicts N_{CCN} very well, with $R_{CCN p/m}$ of 0.90–0.98 410

- 411 (corresponding to a slight underestimation of 2-10%). For the EI-SR scheme,
- 412 hydrophobic POA and a portion of the BC are assumed to be externally mixed while
- the other species (sulfate, nitrate, SOA and aged BC) are assumed to be internal 413
- mixtures. The assumption is physically sound, and the result just implies that the 414
- EI-SR represents well the actual mixing state and compositions of the particles. The 415
- slight underestimation may due to an overestimation of fresh BC caused by the 416
- 417 method (see Section 3.1) that we used to retrieve it. Also, a slightly larger
- underestimation of N_{CCN} for BG case in EI-SR scheme showed in Figure 6 may 418
- suggest that aerosols during clean periods are mostly aged and internal-mixed. 419

420 The INT-SR and INT-BK schemes that assume the aerosol is internally mixed also

predict N_{CCN} reasonably well at lower SS. The prediction is better on background days, 421

422	reflecting the more homogenous aerosol composition in clean conditions. With
423	increasing SS this overestimation became more pronounced, which is likely due to
424	limitations of the AMS measurements. The AMS distributions show that the mass
425	concentration was most impacted by particles with diameters near $\sim 100-400$ nm.
426	Because particles in that size range tended to be more hygroscopic than those with
427	diameters < 100 nm, this leads to an overestimation of κ (underestimation of the
428	critical diameter) and a resulting overestimation of N_{CCN} at high SS. With decreasing
429	SS the critical diameter increased and the deviation using the INT-BK and INT-SR
430	schemes decreased. Detailed explanations about this effect have been given by Wang
431	et al. (2010) and Zhang et al. (2017). Overall, the INT-BK and INT-SR schemes
432	achieve CCN closure within what is deemed here an acceptable overprediction of
433	0-16%. The EXT-BK and EXT-SR schemes underestimated N _{CCN} , with $R_{CCN_p/m}$ of
434	0.66-0.75.

435 Overall, the internal-mixing schemes achieve much better closure than do those assuming external mixtures. Our results suggest that freshly-emitted particles in 436 urban Beijing may experience a quick conversion and mixing with pre-existing 437 438 secondary particles, e.g. converting from externally mixed to internally mixed (or from hydrophobic to hydrophilic, along with a decrease in the volume of POA and BC) 439 as reported previously (Riemer et al., 2004; Aggarwal and Kawamura, 2009; Jimenez 440 441 et al., 2009; Wu et al., 2016; Peng et al., 2016). In summary, under background conditions, the INT-BK scheme achieved the best CCN closure, implying that the 442 INT-BK assumption is likely sufficient to predict CCN in clean continental regions. 443

444 However, in polluted regions, the EI-SR and INT-SR schemes may achieve better445 closure.

446	As mentioned in Section 2.2, because the SP2 measures BC core diameter and
447	not the diameter of the BC-containing particle, the method would overestimate the BC
448	mass concentration of smaller particles but underestimate that of the larger ones. This
449	effect adds uncertainty to the CCN prediction when using the EXT-SR scheme and is
450	evaluated here (Fig. 7). For the evaluation, we predict N_{CCN} with the retrieved fresh
451	BC size distribution only in the EXT-SR scheme, which represents an upper limit of
452	the overestimation of the fresh BC size distribution due to the SP2 measurement.
453	Therefore, the result represents the largest underestimation of N_{CCN} caused by the
454	BC-containing particle effect. Our result shows that the underestimation of N_{CCN} is
455	reduced from 28% to 25% by changing the total BC size distribution to that of just the
456	fresh BC. That means that the overestimation of fresh BC due to the BC-containing
457	particle effect in the SP2 measurements would lead to a maximum underestimation of
458	3% of N_{CCN} . The minimal uncertainty contributed by uncertainty in the BC size
459	distribution could be explained by the small fractional contribution of BC to the total
460	particle concentration. In conclusion, such an effect is quite small or negligible
461	compared to the overall large underestimation of N_{CCN} with the EXT-SR assumption.

462 **4.4 Performance of the five schemes at different times of the day**

463 To investigate the performance of the five schemes at different times of the day, 464 the diurnal variations in the $R_{CCN_p/m}$ (SS = 0.23%) derived by the schemes are shown

465	in Fig. 8. In general, the INT-BK, INT-SR, and EI-SR schemes can predict N_{CCN}
466	well during all periods of the day under polluted or background conditions. $R_{CCN_p/m}$
467	values are within the acceptable $\pm 20\%$ uncertainty range (Wang et al., 2010; Zhang
468	et al., 2017). Compared with other periods, the predicted N_{CCN} during the morning
469	and evening rush hour periods showed the most sensitivity to the different assumption
470	schemes, especially on clean days (Fig. 8b). For example, the $R_{CCN_p/m}$ derived using
471	the INT-SR schemes reaches values up to >1.2, and the $R_{CCN_p/m}$ obtained using the
472	EXT-BK scheme decreased to a minimum value of ~0.5. The INT-SR, INT-BK and
473	EI-SR assumptions overestimate N_{CCN} for the evening rush hour period by up to
474	\sim 20%. This may be because most freshly emitted POA and BC particles during
475	evening traffic hours are hydrophobic and do not contribute to the $N_{\mbox{\scriptsize CCN}}.$ In addition,
476	for EIS assumption, a portion of BC is assumed aged and internally-mixed with
477	sulfate, nitrate and SOA, as may reduce the actual fraction of fresh BC during rush
478	hour period and thereby lead to an overestimation of N_{CCN} .
479	Use of the EXT-BK or EXT-SR assumption for the polluted case resulted in a
480	predicted N_{CCN} that was underestimated by ~30-40% at night (0000–0600 LT).
481	Expectedly, the prediction using the two schemes improved during the daytime and
482	evening rush hours, e.g., the $R_{CCN_p/m}$ changed from about 0.6 to 0.8. This is likely
483	associated with heavy urban traffic emissions/residential cooking sources during the
484	daytime that lead to more externally-mixed particles under polluted conditions; while,
485	at night, the particles are less influenced by those local primary sources (Zhao. et al.,

486 2017). Wang et al. (2017) showed that the probability density function of κ during

rush hour has a bimodal distribution and a hydrophobic mode from locally-emitted particles. This also leads to reasonably accurate estimates of N_{CCN} during nighttime with larger error during the daytime when using the internal mixing assumptions (INT-BK, INT-SR and EI-SR) for polluted cases (Fig. 8).

491 **4.5 Impact of mixing state and organic volume fraction on predicted N**_{CCN} and

492 their variation with aerosol aging

493 To further examine the sensitivity of predicted N_{CCN} to the particle mixing state and organic volume fraction with the aging of organic particles, the relative deviation 494 between N_{CCN} predicted assuming internal and external mixtures as a function of κ_{org} 495 was calculated, with the results shown in Fig. 9. The schemes that assume internal and 496 497 external mixtures use bulk composition of organics, sulfate, and nitrate, which simplifies the analysis and interpretation of the results. For the data collected 498 throughout the field campaign, the organic volume fraction is categorized as <50%, 499 50-60%, and >60%. The deviation between the concentrations predicted assuming 500 internal and external mixtures is calculated as [(N_{CCN}INT-BK - N_{CCN}EXT-BK) 501 (N_{CCN}.EXT-BK)⁻¹]. The result shows that the relative deviation increased as the 502 503 organic volume fraction increased. For organic volume fractions less than 50% the maximum difference can only reach up to 20% (SS=0.76%). This is consistent with 504 previous studies that reported differences less than 20% when $x_{org} < 30\%$ 505 506 (Sotiropoulou et al., 2006; Wang et al., 2010). The maximum deviation approaches to 100% for x_{org} of >60% at SS = 0.76%. Overall, the deviation is largest when the 507

organics are less or non-hygroscopic, i.e., when $\kappa_{org} < 0.05$. The deviation decreased rapidly as κ_{org} increased to 0.05 in all cases. For κ_{org} of 0.1 the differences were less than 20%, even with high organic fractions. Moreover, differences were 10% or less for κ_{org} of 0.15, suggesting that the mixing state plays a minor role when κ_{org} exceeds 0.1.

513 **5 Conclusions**

514	In this study, we have investigated the importance of aerosol chemical
515	composition and mixing state on CCN activity based on measurements made during a
516	field campaign carried out in Beijing in the winter of 2016. The N_{CCN} was predicted
517	by applying κ -Köhler theory and using five schemes that assume different mixing
518	state and chemical composition combinations.
519	We show that there is a significant impact of the different assumptions on CCN
520	prediction, with $R_{CCN_p/m}$ ranging from 0.66 to 1.16. The best estimates of N _{CCN} under
521	both background and polluted conditions were obtained when using the EI-SR scheme,
522	with a resulting $R_{CCN_p/m}$ of 0.90–0.98. Under background conditions, the INT-BK
523	scheme also provided reasonable estimates, with $R_{CCN_p/m}$ ranging from 1.00–1.16.
524	This suggests that the INT-BK assumption is likely sufficient to predict CCN in clean
525	continental regions. On polluted days, the EI-SR and INT-SR schemes are believed to
526	achieve better closure than the INT-BK scheme due to the heterogeneity in particle
527	composition across different sizes. The improved closure obtained using the EI-SR
528	and INT-SR assumptions highlights the importance of knowing the size-resolved

529	chemical composition for CCN prediction in polluted regions. The EXT-SR and
530	EXT-BK schemes markedly underestimate N_{CCN} on both polluted and clean days,
531	with an $R_{CCN_p/m}$ of 0.66–0.75. The diurnal variations in $R_{CCN_p/m}$ show that the
532	predicted N_{CCN} during the evening rush hour period is most sensitive to the mixing
533	state assumptions. The $R_{CCN_p/m}$ ranged from ~0.5 to ~1.2, reflecting the impact from
534	evening traffic and cooking sources (both with large amounts of hydrophobic POA).
535	But we also find that the particle mixing state plays a minor role when κ_{org} exceeds
536	0.1, even with a high organic fraction.
537	Acknowledgements. This work was funded by the NSFC research project (41675141
538	and 91544217), the fundamental Research Funds for the Central Universities, the
539	National Basic Research Program of China '973' (2013CB955800), the NSCF-TAMU
540	Collaborative Research Grant Program (4141101031), and the Natural Science
541	Foundation (NSF) (AGS1534670). We thank all participants of the field campaign for
542	their tireless work and cooperation.

544 **References**

- 545 Aggarwal, S. G., and Kawamura, K.: Carbonaceous and inorganic composition in long-range
- 546 transported aerosols over northern Japan: Implication for aging of water-soluble organic
- 547 fraction, Atmos. Environ., 43, 2532–2540, doi:10.1016/j.atmosenv.2009.02.032, 2009.
- Albrecht, B. A.: Aerosols, cloud microphysics, and fractional cloudiness, Science, 245, 1227–1230,
 1989.
- 550 Andreae, M. O., and Rosenfeld, D.: Aerosol-cloud-precipitation interactions. Part 1. The nature
- and sources of cloud-active aerosols, Earth-Science Reviews, 89, 13–41,
- 552 doi:10.1016/j.earscirev.2008.03.001, 2008.
- 553 Aiken, A. C., Salcedo, D., Cubison, M. J., Huffman, J. A., DeCarlo, P. F., Ulbrich, I. M.,
- 554 Docherty, K. S., Sueper, D., Kimmel, J. R., Worsnop, D. R., Trimborn, A., Northway, M.,
- 555 Stone, E. A., Schauer, J. J., Volkamer, R. M., Fortner, E., de Foy, B., Wang, J., Laskin, A.,
- 556 Shutthanandan, V., Zheng, J., Zhang, R., Gaffney, J., Marley, N. A., Paredes-Miranda, G.,
- 557 Arnott, W. P., Molina, L. T., Sosa, G., and Jimenez, J. L.: Mexico City aerosol analysis
- during MILAGRO using high resolution aerosol mass spectrometry at the urban supersite
- 559 (T0) Part 1: Fine particle composition and organic source apportionment, Atmos. Chem.
- 560 Phys., 9, 6633–6653, doi:10.5194/acp-9-6633-2009, 2009.
- Bhattu, D., and Tripathi, S. N.: CCN closure study: Effects of aerosol chemical composition and
 mixing state, J. Geophys. Res. Atmos., 120, 766–783, doi:10.1002/2014jd021978, 2015.
- 563 Broekhuizen, K., Chang, R. Y. W., Leaitch, W. R., Li, S. M., and Abbatt, J. P. D.: Closure between
- 564 measured and modeled cloud condensation nuclei (CCN) using size-resolved aerosol
- 565 compositions in downtown Toronto, Atmos. Chem. Phys., 6, 2513–2524,
- 566 10.5194/acp-6-2513-2006, 2006.
- Chang, R. Y. W., Liu, P. S. K., Leaitch, W. R., and Abbatt, J. P. D.: Comparison between measured
 and predicted CCN concentrations at Egbert, Ontario: Focus on the organic aerosol fraction at a
- 569 semirural site, Atmos. Environ., 41, 8172–8182, 2007.
- 570 Charlson, R. J., Schwartz, S. E., Hales, J. M., Cess, R. D., Coakley, J. A., Jr., Hansen, J. E., and
- 571 Hofmann, D. J.: Climate forcing by anthropogenic aerosols, Science, 255, 423+, 1992.
- 572 Che, H. C., Zhang, X. Y., Wang, Y. Q., Zhang, L., Shen, X. J., Zhang, Y. M., Ma, Q. L., Sun, J. Y.,
- 573 Zhang, Y. W., and Wang, T. T.: Characterization and parameterization of aerosol cloud
- 574 condensation nuclei activation under different pollution conditions, Sci. Rep, 6,
- 575 doi:10.1038/srep24497, 2016.

- 576 Cross, E. S., Onasch, T. B., Canagaratna, M., Jayne, J. T., Kimmel, J., Yu, X. Y., Alexander, M. L.,
- 577 Worsnop, D. R., and Davidovits, P.: Single particle characterization using a light scattering
- 578 module coupled to a time of flight aerosol mass spectrometer, Atmos. Chem. Phys., 9,
- 579 7769–7793, doi:10.5194/acp-9-7769-2009, 2009.
- 580 Cheng, Y. F., Su, H., Rose, D., Gunthe, S. S., Berghof, M., Wehner, B., Achtert, P., Nowak, A.,
- 581 Takegawa, N., Kondo, Y., Shiraiwa, M., Gong, Y. G., Shao, M., Hu, M., Zhu, T., Zhang, Y. H.,
- 582 Carmichael, G. R., Wiedensohler, A., Andreae, M. O., and Pöschl, U.: Size-resolved
- 583 measurement of the mixing state of soot in the megacity Beijing, China: diurnal cycle, aging
- and parameterization, Atmospheric Chemistry and Physics, 12, 4477-4491,
- 585 10.5194/acp-12-4477-2012, 2012.
- 586 Dall'Osto, M., Harrison, R. M., Coe, H., Williams, P. I., and Allan, J. D.: Real time chemical
- characterization of local and regional nitrate aerosols, Atmos. Chem. Phys., 9, 3709–3720,
 10.5194/acp-9-3709-2009, 2009.
- 589 DeCarlo, P. F., Kimmel, J. R., Trimborn, A., et al.: Field-deployable, high-resolution,
- time-of-flight aerosol mass spectrometer, Anal. Chem., 78, 8281–8289, 2006.
- Deng, Z. Z., Zhao, C. S., Ma, N., Ran, L., Zhou, G. Q., Lu, D. R., and Zhou, X. J.: An examination
 of parameterizations for the CCN number concentration based on in situ measurements of
 aerosol activation properties in the North China Plain, Atmos. Chem. Phys., 13, 6227–6237,
- 594 10.5194/acp-13-6227-2013, 2013.
- 595 Dusek, U., Frank, G. P., Hildebrandt, L., et al.: Size matters more than chemistry for cloud
- nucleating ability of aerosol particles, Science, 312, 1375–1378, 2006.
- 597 Dzepina, K., Volkamer, R. M., Madronich, S., Tulet, P., Ulbrich, I. M., Zhang, Q., Cappa, C. D.,
- 598 Ziemann, P. J., and Jimenez, J. L.: Evaluation of recently proposed secondary organic aerosol
- 599 models for a case study in Mexico City, Atmos. Chem. Phys., 9, 5681–5709,
- 600 doi:10.5194/acp-9-5681-2009, 2009.
- 601 Ervens, B., Cubison, M., Andrews, E., et al.: Prediction of cloud condensation nucleus number
- 602 concentration using measurements of aerosol size distributions and composition and light
- scattering enhancement due to humidity, J. Geophys. Res. Atmos., 112, D10S32,
- 604 doi:10.1029/2006JD007426, 2007.
- Gunthe, S. S., King, S. M., Rose, D., Chen, Q., Roldin, P., Farmer, D. K., Jimenez, J. L., Artaxo, P.,
- Andreae, M. O., Martin, S. T., and Pöschl, U.: Cloud condensation nuclei in pristine tropical
- 607 rainforest air of Amazonia: size resolved measurements and modeling of atmospheric aerosol
- 608 composition and CCN activity, Atmos. Chem. Phys., 9, 7551–7575,
- 609 doi:10.5194/acp-9-7551-2009, 2009.

- 610 Gunthe, S. S., Rose, D., Su, H., Garland, R. M., Achtert, P., Nowak, A., Wiedensohler, A., Kuwata,
- 611 M., Takegawa, N., Kondo, Y., Hu, M., Shao, M., Zhu, T., Andreae, M. O., and Pöschl, U.: Cloud
- 612 condensation nuclei (CCN) from fresh and aged air pollution in the megacity region of Beijing,
- 613 Atmos. Chem. Phys., 11, 11023–11039, doi:10.5194/acp-11-11023-2011, 2011.
- 614 Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M., Zeng, L.,
- Molina, M. J., and Zhang, R.: Elucidating severe urban haze formation in China, P. Natl. Acad.
- 616 Sci. USA, 111, 17373–17378, doi:10.1073/pnas.1419604111, 2014.
- 617 Gysel, M., Crosier, J., Topping, D. O., Whitehead, J. D., Bower, K. N., Cubison, M. J., Williams, P.
- 618 I., Flynn, M. J., McFiggans, G. B., and Coe, H.: Closure study between chemical composition
- and hygroscopic growth of aerosol particles during TORCH2, Atmos. Chem. Phys., 7,
- 620 6131–6144, doi:10.5194/acp-7-6131-2007, 2007.
- Jimenez, J. L., Canagaratna, M. R., et al.: Evolution of organic aerosols in the atmosphere,
 Science, 326, 1525–1529, 2009.
- 623 Kawana, K., Nakayama, T., and Mochida, M.: Hygroscopicity and CCN activity of atmospheric
- aerosol particles and their relation to organics: Characteristics of urban aerosols in Nagoya,
- 625 Japan, J. Geophys. Res. Atmos., 121, 4100–4121, doi:10.1002/2015jd023213, 2016.
- Lance, S., Medina, J., Smith, J., and Nenes, A.: Mapping the operation of the DMT continuous
 flow CCN counter, Aerosol Sci. Technol., 40, 242–254, 2006.
- 628 Li, Y., Zhang, F., Li, Z., Sun, L., Wang, Z., Li, P., Sun, Y., Ren, J., Wang, Y., Cribb, M., and Yuan,
- 629 C.: Influences of aerosol physiochemical properties and new particle formation on CCN activity
- from observation at a suburban site of China, Atmos. Res., 188, 80–89,
- 631 doi:10.1016/j.atmosres.2017.01.009, 2017.
- Li, Z., F. Niu, J. Fan, Y. Liu, D. Rosenfeld, and Y. Ding.: The long-term impacts of aerosols on the
 vertical development of clouds and precipitation, Nature Geosci. 4, doi: 10.1038/NGEO1313,
 2011.
- Liu, H. J., Zhao, C. S., Nekat, B., Ma, N., Wiedensohler, A., van Pinxteren, D., Spindler, G.,
- 636 Müller, K., and Herrmann, H.: Aerosol hygroscopicity derived from size-segregated chemical
- 637 composition and its parameterization in the North China Plain, Atmos. Chem. Phys., 14,
- 638 2525–2539, doi:10.5194/acp-14-2525-2014, 2014.
- Ma, Y., Brooks, S. D., Vidaurre, G., Khalizov, A. F., Wang, L., and Zhang, R..: Rapid modification
 of cloud-nucleating ability of aerosols by biogenic emissions, Geophys. Res. Lett., 40(23),
- 641 6293–6297, 2013.

642	McFiggans, G., Artaxo, P., Baltensperger, U., Coe, H., Facchini, M. C., Feingold, G., Fuzzi, S.,
643	Gysel, M., Laaksonen, A., Lohmann, U., Mentel, T. F., Murphy, D. M., O'Dowd, C. D., Snider,
644	J. R., and Weingartner, E.: The effect of physical and chemical aerosol properties on warm
645	cloud droplet activation, Atmos. Chem. Phys., 6, 2593–2649, doi:10.5194/acp-6-2593-2006,
646	2006.
647	Medina, J., Nenes, A., Sotiropoulou, R. E. P., Cottrell, L. D., Ziemba, L. D., Beckman, P. J., and
648	Griffin, R. J.: Cloud condensation nuclei closure during the International Consortium for
649	Atmospheric Research on Transport and Transformation 2004 campaign: Effects of size
650	resolved composition, J. Geophys. Res. Atmos., 112, D10S31, doi:10.1029/2006JD007588,
651	2007.
652	Mei, F., Setyan, A., Zhang, Q., and Wang, J.: CCN activity of organic aerosols observed
653	downwind of urban emissions during CARES, Atmos. Chem. Phys., 13, 12155-12169,
654	doi:10.5194/acp-13-12155-2013, 2013a.
655	Mei, F., Hayes, P. L., Ortega, A. M., Taylor, J. W., Allan, J. D., Gilman, J. B., Kuster, W. C., de
656	Gouw, J. A., Jimenez, J. L., and Wang, J.: Droplet activation properties of organic aerosols
657	observed at an urban site during CalNex-LA, J. Geophys. Res., 118, 2903-2917,
658	10.1002/jgrd.50285, 2013b.
659	Meng, J. W., Yeung, M. C., Li, Y. J., Lee, B. Y. L., and Chan, C. K.: Size-resolved cloud
660	condensation nuclei (CCN) activity and closure analysis at the HKUST Supersite in Hong Kong,
661	Atmos. Chem. Phys., 14, 10267–10282, doi:10.5194/acp-14-10267-2014, 2014.
662	Moore, R. H., Nenes, A., and Medina, J.: Scanning mobility CCN analysis-A method for fast
663	measurements of size-resolved CCN distributions and activation kinetics, Aerosol Sci. Technol.,
664	44, 861–871, doi:10.1080/02786826.2010.498715, 2010.
665	Moore, R. H., Cerully, K., Bahreini, R., Brock, C. A., Middlebrook, A. M., and Nenes, A.:
666	Hygroscopicity and composition of California CCN during summer 2010, J. Geophys. Res.
667	Atmos., 117, D00V12, doi:10.1029/2011JD017352, 2012.
668	Paatero, P., and U. Tapper (1994), Positive matrix factorization: A non-negative factormodel with
669	optimal utilization of error estimates of data values, Environmetrics, 5, 111-126.
670	Peng, J., Hu, M., Guo, S., Du, Z., Zheng, J., & Shang, D., et al. (2016). Markedly enhanced
671	absorption and direct radiative forcing of black carbon under polluted urban environments.
672	Proceedings of the National Academy of Sciences of the United States of America, 113(16),
673	4266.
674	Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth

- and cloud condensation nucleus activity, Atmos. Chem. Phys., 7, 1961–1971,
- 676 doi:10.5194/acp-7-1961-2007, 2007.
- Riemer, N., Vogel, H., and Vogel, B.: Soot aging time scales in polluted regions during day and
 night, Atmos. Chem. Phys., 4, 1885–1893, doi:10.5194/acp-4-1885-2004, 2004.
- 679 Rose, D., Gunthe, S. S., Mikhailov, E., Frank, G. P., Dusek, U., Andreae, M. O., and Pöschl, U.:
- 680 Calibration and measurement uncertainties of a continuous-flow cloud condensation nuclei
- 681 counter (DMT-CCNC): CCN activation of ammonium sulfate and sodium chloride aerosol
- particles in theory and experiment, Atmos. Chem. Phys., 8, 1153–1179,
- 683 doi:10.5194/acp-8-1153-2008, 2008.
- 684 Rose, D., Nowak, A., Achtert, P., Wiedensohler, A., Hu, M., Shao, M., Zhang, Y., Andreae, M. O.,
- and Pöschl, U.: Cloud condensation nuclei in polluted air and biomass burning smoke near the
- 686 mega-city Guangzhou, China. Part 1: Size-resolved measurements and implications for the
- 687 modeling of aerosol particle hygroscopicity and CCN activity, Atmos. Chem. Phys., 10,
- 688 3365–3383, doi:10.5194/acp-10-3365-2010, 2010.
- 689 Rosenfeld, D., Lohmann, U., Raga, G. B., O'Dowd, C. D., Kulmala, M., Fuzzi, S., Reissell, A.,
- and Andreae, M. O.: Flood or drought: How do aerosols affect precipitation?, Science, 321,
 doi:10.1126/science.1160606, 2008.
- 692 Sotiropoulou, R.-E. P., Medina, J., and Nenes, A.: CCN predictions: Is theory sufficient for
- assessments of the indirect effect?, Geophys. Res. Lett., 33, doi:10.1029/2005gl025148, 2006.
- 694 Stokes, R. H., and Robinson, R. A.: Interactions in aqueous nonelectrolyte solutions. I.
- 695 Solute-solvent equilibria, J. Phys. Chem., 70, 2126–2130, 1966.
- 696 Sun, J., Zhang, Q., Canagaratna, M. R., Zhang, Y., Ng, N. L., Sun, Y., Jayne, J. T., Zhang, X.,
- Zhang, X., and Worsnop, D. R.: Highly time- and size-resolved characterization of submicron
 aerosol particles in Beijing using an Aerodyne Aerosol Mass Spectrometer, Atmos. Environ., 44,
- 699 131–140, doi:10.1016/j.atmosenv.2009.03.020, 2010.
- Sun, Y. L., Wang, Z. F., Fu, P. Q., Yang, T., Jiang, Q., Dong, H. B., Li, J., and Jia, J. J.: Aerosol
- composition, sources and processes during wintertime in Beijing, China, Atmos. Chem. Phys.,
- 702 13, 4577–4592, doi:10.5194/acp-13-4577-2013, 2013.
- 703 Sun, Y. L., Wang, Z. F., Du, W., Zhang, Q., Wang, Q. Q., Fu, P. Q., Pan, X. L., Li, J., Jayne, J., and
- Worsnop, D. R.: Long-term real-time measurements of aerosol particle composition in Beijing,
- 705 China: Seasonal variations, meteorological effects, and source analysis, Atmos. Chem. Phys., 15,
- 706 10149–10165, doi:10.5194/acp-15-10149-2015, 2015.

- 707 Sun, Y., Chen, C., Zhang, Y., Xu, W., Zhou, L., Cheng, X., Zheng, H., Ji, D., Li, J., Tang, X., Fu, P.,
- and Wang, Z.: Rapid formation and evolution of an extreme haze episode in Northern China
 during winter 2015, Sci. Rep., 6, doi:10.1038/srep27151, 2016.
- 710 Textor, C., Schulz, M., Guibert, S., Kinne, S., Balkanski, Y., Bauer, S., Berntsen, T., Berglen, T.,
- 711 Boucher, O., Chin, M., Dentener, F., Diehl, T., Easter, R., Feichter, H., Fillmore, D., Ghan, S.,
- Ginoux, P., Gong, S., Grini, A., Hendricks, J., Horowitz, L., Huang, P., Isaksen, I., Iversen, I.,
- 713 Kloster, S., Koch, D., Kirkevåg, A., Kristjansson, J. E., Krol, M., Lauer, A., Lamarque, J. F., Liu,
- X., Montanaro, V., Myhre, G., Penner, J., Pitari, G., Reddy, S., Seland, Ø., Stier, P., Takemura,
- T., and Tie, X.: Analysis and quantification of the diversities of aerosol life cycles within
- 716 AeroCom, Atmos. Chem. Phys., 6, 1777–1813, doi:10.5194/acp-6-1777-2006, 2006.
- 717 Twomey, S.: The influence of pollution on the shortwave albedo of clouds, J. Atmos. Sci., 34,

- 719 Wang, J., Flagan, R. C., and Seinfeld, J. H.: A differential mobility analyzer (DMA) system for
- submicron aerosol measurements at ambient relative humidity, Aerosol Sci. Technol., 37, 46–52,
 2003.
- Wang, J., Cubison, M. J., Aiken, A. C., Jimenez, J. L., and Collins, D. R.: The importance of
 aerosol mixing state and size-resolved composition on CCN concentration and the variation of
 the importance with atmospheric aging of aerosols, Atmos. Chem. Phys., 10, 7267–7283,
 dei:10.5104/sec.10.7267.2010.2010.
- doi:10.5194/acp-10-7267-2010, 2010.
- 726 Wang, G., Zhang, R., Gomez, M. E., Yang, L., Zamora, M. L., Hu, M., ... & Li, J. (2016).
- Persistent sulfate formation from London Fog to Chinese haze. Proceedings of the National
- Academy of Sciences, 113(48), 13630-13635.
- Wang, Y., Zhang, F., Li, Z., Tan, H., Xu, H., Ren, J., Zhao, J., Du, W., and Sun, Y.: Enhanced
- hydrophobicity and volatility of submicron aerosols under severe emission control conditions in
 Beijing, Atmos. Chem. Phys., 17, 5239–5251, doi:10.5194/acp-17-5239-2017, 2017.
- 732 Wiedensohler, A., Cheng, Y. F., Nowak, A., Wehner, B., Achtert, P., Berghof, M., Birmili, W., Wu,
- 733 Z. J., Hu, M., Zhu, T., Takegawa, N., Kita, K., Kondo, Y., Lou, S. R., Hofzumahaus, A., Holland,
- F., Wahner, A., Gunthe, S. S., Rose, D., Su, H., and Pöschl, U.: Rapid aerosol particle growth
- and increase of cloud condensation nucleus activity by secondary aerosol formation and
- condensation: A case study for regional air pollution in northeastern China, J. Geophys. Res.
- 737 Atmos., 114, D00G08, doi:10.1029/2008JD010884, 2009.
- 738 Wu, Y., Wang, X., Tao, J., Huang, R., Tian, P., Cao, J., Zhang, L., Ho, K.-F., Han, Z., and Zhang,
- R.: Size distribution and source of black carbon aerosol in urban Beijing during winter haze
- r40 episodes, Atmos. Chem. Phys., 17, 7965–7975, doi:10.5194/acp-17-7965-2017, 2017.

^{718 1149–1152,} doi:10.1175/1520-0469(1977)034(1149: TIOPOT)2.0.CO;2, 1977.

- 741 Wu, Z. J., Zheng, J., Shang, D. J., Du, Z. F., Wu, Y. S., Zeng, L. M., Wiedensohler, A., and Hu, M.:
- Particle hygroscopicity and its link to chemical composition in the urban atmosphere of Beijing,
- China, during summertime, Atmos. Chem. Phys., 16, 1123–1138,
- 744 doi:10.5194/acp-16-1123-2016, 2016.
- 745 Yum, S. S., Hudson, J. G., Song, K. Y., and Choi, B. C.: Springtime cloud condensation nuclei
- concentrations on the west coast of Korea, Geophys. Res. Lett., 32, L09814,
- 747 doi:10.1029/2005GL022641, 2005.
- Yum, S. S., Roberts, G., Kim, J. H., Song, K. Y., and Kim, D. Y.: Submicron aerosol size
- distributions and cloud condensation nuclei concentrations measured at Gosan, Korea, during
- the Atmospheric Brown Clouds East Asian Regional Experiment 2005, J. Geophys. Res. Atmos.,

•

- 751 112, D22S32, doi:10.1029/2006JD008212, 2007.
- Zdanovskii, B.: Novyi Metod Rascheta Rastvorimostei Elektrolitovv Mnogokomponentnykh
 Sistema, Zh. Fiz. Khim+, 22, 1478–1495, 1948.
- Zhang, F., Li, Y., Li, Z., Sun, L., Li, R., Zhao, C., Wang, P., Sun, Y., Liu, X., Li, J., Li, P., Ren, G.,
 and Fan, T.: Aerosol hygroscopicity and cloud condensation nuclei activity during the AC³Exp
- campaign: Implications for cloud condensation nuclei parameterization, Atmos. Chem. Phys.,
 14, 13423–13437, doi:10.5194/acp-14-13423-2014, 2014.
- 758 Zhang, F., Li, Z., Li, Y., Sun, Y., Wang, Z., Li, P., Sun, L., Wang, P., Cribb, M., Zhao, C., Fan, T.,
- Yang, X., and Wang, Q.: Impacts of organic aerosols and its oxidation level on CCN activity
- from measurement at a suburban site in China, Atmos. Chem. Phys., 16, 5413–5425,
- 761 doi:10.5194/acp-16-5413-2016, 2016.
- Zhang F., Wang Y., Peng J., Ren J., Zhang R., Sun Y., Don Collin., Yang X., Li Z.: Uncertainty
 in predicting CCN activity of aged and primary aerosols. Journal of Geophysical
 Research: Atmospheres. 10.1002/2017JD027058, 2017.
- 765 Zhang, Z., Engling, G., Lin, C.-Y., Chou, C. C. K., Lung, S.-C. C., Chang, S.-Y., Fan, S., Chan,
- 766 C.-Y., and Zhang, Y.-H.: Chemical speciation, transport and contribution of biomass burning
- smoke to ambient aerosol in Guangzhou, a mega city of China, Atmos. Environ., 44, 3187–3195,
- 768 doi:10.1016/j.atmosenv.2010.05.024, 2010.
- 769 Zhao, J., Du, W., Zhang, Y., Wang, Q., Chen, C., Xu, W., Han, T., Wang, Y., Fu, P., Wang, Z., Li,
- Z., and Sun, Y.: Insights into aerosol chemistry during the 2015 China Victory Day parade:
- results from simultaneous measurements at ground level and 260 m in Beijing, Atmos. Chem.
- Phys., 17, 3215–3232, doi:10.5194/acp-17-3215-2017, 2017.
- 773 Zheng, G. J., Duan, F. K., Su, H., Ma, Y. L., Cheng, Y., Zheng, B., Zhang, Q., Huang, T., Kimoto,

1/4 I., Chang, D., Poschi, U., Cheng, Y. F., and He, K. B.: Exploring the severe	winter naze	; 1n
--	-------------	------

- Beijing: the impact of synoptic weather, regional transport and heterogeneous reactions, Atmos.
- 776 Chem. Phys., 15, 2969–2983, doi:10.5194/acp-15-2969-2015, 2015.

777	Zhang, Q., Worsnop, D. R., Canagaratna, M. R., and Jimenez, J. L.: Hydrocarbon-like and
778	oxygenated organic aerosols in Pittsburgh: insights into sources and processes of organic
779	aerosols, Atmos. Chem. Phys., 5, 3289-3311, doi:10.5194/acp-5-3289- 2005, 2005.

- /04

795 Figures



Figure 1 Schematic representation of the five different schemes: (a) INT-BK, (b) 797 EXT-BK, (c) EI-SR, (d) INT-SR, and (e) EXT-SR. And the BC size distribution (f) 798 used in this study. The fresh and aged BC size distribution are retrieved from the total 799 BC size distribution measured by the SP2 (Wu et al., 2017) and the dependence of the 800 fraction of internally mixed soot (F_{in}) on particle diameter (D_p) observed in urban 801 Beijing (Cheng et al., 2012). The total BC size distribution is used in the INT-SR and 802 803 EXT-SR schemes, and the aged and fresh BC distributions are used in the EI-SR scheme. In the EI-SR scheme, some BC particles are assumed to already be aged and 804 thus internally-mixed with sulfate, nitrate and SOA, and some of them together with 805 POA are freshly emitted and assumed not yet aged/coated by other species 806

807 (externally-mixed).



Figure 2. Diurnal variations in aerosol properties at the IAP site during the APHH
field experiment, including the particle number size distribution measured by the

- 813 SMPS under (a) background (BG) and (b) polluted (POL) conditions; (c) mean
- 814 particle number size distribution measured by the SMPS during three periods
- 815 (0000–0200 LT, 1200–1400 LT, and 1700–2000 LT) under BG and POL conditions;
- bulk chemical component mass concentrations (NO₃, POA, SOA, and BC) and f_{44}
- 817 made under (d) BG and (e) POL conditions.
- 818



Figure 3. (a) Averaged fitted CCN efficiency spectra during the nighttime period

821 (0000–0200 LT, dashed lines), the noontime period (1200–1400 LT, dotted lines) and

the evening rush hour period (1700–2000 LT, solid lines) for different diameters (60,

823 100, 150, and 200 nm); (b) the heterogeneity of aerosol particles (σ_a/D_a) derived from

824 Equation (7) during the three selected periods.



Figure 4. Top: Retrieved mean critical activation diameters at SS = 0.12, 0.14, 0.23, 0.14, 0.23

829 0.40, and 0.76% under background (BG) and polluted (POL) conditions. The box

plots show mean critical activation diameters at the 25th, 50th, and 75th percentiles.

831 Bottom: Difference in the mean critical activation diameter between BG and POL

832 cases.

833



836 Figure 5. Mean maximum active fractions (MAFs) of CCN activation spectra under

837 polluted (POL) and background (BG) conditions during the three periods, i.e.,

838 0000–0200 LT, 1200–1400 LT, and 1700–2000 LT. The MAF of pure (NH₄)₂SO₄

839 particles at the different SS levels (magenta line) is also plotted.



842

843 + INT-BK Internal mixture, bulk composition



- 845 * EXT-BK External mixture, bulk composition
- 846 Description EXT-SR External mixture, size-resolved composition
- 848 Figure 6. Predicted N_{CCN} as a function of measured N_{CCN} using the five assumptions
- (colored symbols) at three supersaturation levels (0.23, 0.40, and 0.76%) under
- 850 polluted (POL) and background (BG) conditions. The numbers in parentheses are the
- slope (first number) and the correlation coefficient (second number).



Figure 7. Predicted N_{CCN} as a function of measured N_{CCN} using the EXT-SR assumption (colored symbols) at S=0.76%. The pink and blue circles denote the results predicted by using total and fresh BC size distributions, respectively. The numbers in parentheses are the slope (first number) and the correlation coefficient (second number).

861

862

863



- 866 + INT-BK Internal mixture, bulk composition
- 867 **o** INT-SR Internal mixture, size-resolved composition
- 868 * EXT-BK External mixture, bulk composition
- 869 EXT-SR External mixture, size-resolved composition
- 870 **v** EI-SR External mixture, POA and BC externally mixed, size-resolved composition
- Figure 8. Diurnal variations in the ratio of predicted-to-measured N_{CCN} at a
- supersaturation level of 0.23% under background (BG) and polluted (POL)
- 873 conditions.



Figure 9. Relative deviations between N_{CCN} predicted under the assumptions of internal (INT-BK) and external (EXT-BK) mixtures [(N_{CCN}.INT-BK - N_{CCN}.EXT-BK) (N_{CCN}.EXT-BK)⁻¹] as a function of κ_{org} when organic volume fractions of <50 (a), 50-60 (b), >60% (c) and all observed data points (d). The solid with different colors represent different supersaturation levels. The different colors denote the different organic fractions.

891