



1 **Increasing persistent hazes in Beijing: potential impacts of weakening East Asian Winter**

2 **Monsoons associated with northwestern Pacific SST trend since 1900.**

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11 **Abstract:**

12 This paper analyzes the variations of persistent haze days in boreal winter in Beijing and
13 demonstrates a significant increasing trend with increasing probability of the persistent haze
14 events (PHEs) during 1980-2016. The atmospheric conditions favorable for the increasing
15 PHEs include reduced surface meridional winds and weakened East Asian Trough in the
16 mid-troposphere. These conditions indicate a weakening East Asian Winter Monsoon (EAWM)
17 system, which is then found to be associated with an anomalous warm and high-pressure
18 system in the mid-lower troposphere over the northwestern Pacific. We proposed a practical
19 circulation indicator using the anomalous southerly winds at 850 hPa over North China. The
20 increasing occurrence of the persistent haze days in Beijing is closely related with increasing
21 frequencies of unusual persistent southerly episodes in recent decades. Over the past
22 centennial period 1900-2009, the anomalous meridional winds at 850 hPa in North China
23 were positively correlated with the sea surface temperature anomalies over the northwestern
24 Pacific. Based on the present results, we outline an observation-based mechanism
25 highlighting the role of large-scale climate warming on the increasing trend of PHEs in
26 Beijing.

27

28 **1 Introduction**

29 In the past decades, pollutant emissions have considerably increased in China due to the rapid



30 economic development (Guo et al., 2011; Zhou et al., 2010). The capital city Beijing has
31 consequently encountered with an increasing frequency of severe hazes especially in the
32 winter (Niu et al., 2010; Ding and Liu, 2014; Chen and Wang, 2015; Wang et al., 2015).
33 Notable is the increasing tendency of persistence haze events (PHEs) during the past decades
34 (Zhang et al., 2015; Wu et al., 2017). Severe haze with a high concentration of fine particles
35 such as PM_{2.5} not only leads to a sharp decrease in visibility causing traffic hazards and
36 disruptions and hence affecting economic activities (Chen and Wang, 2015; Li et al., 2016;
37 Huang et al., 2014), but also induces serious health problems dealing with respiratory
38 illnesses, heart disease, premature death and cancers (Pope and Dockery, 2006; Wang and
39 Mauzerall, 2006; Hou et al., 2012; Sun et al., 2013; Xie et al., 2014; Gao et al., 2015). PHEs
40 would aggravate these serious consequences. During 14-25 January 2013, eastern China was
41 hit by a persistent severe haze event affecting about 800 million people, while Beijing reached
42 its highest level of air pollution in record and announced the first orange haze alert (Ding and
43 Liu, 2014; Zhang et al., 2014). In this period, about 200 cases of premature death and
44 thousands of cases of hospital admissions for respiratory, cardiovascular and asthma
45 bronchitis diseases were found to be associated with the high level of PM_{2.5} (Xu et al., 2013).
46 Correspondingly, this event caused about 489 million RMB health-related economic losses.
47 Thus, detailed studies on the characteristics and underlying reasons of the increasing PHEs
48 around Beijing are urgently needed.

49 Many studies suggested that the increased emissions of pollutants into the atmosphere due to
50 rapid economic development and urbanization in China were the main cause of the increasing
51 haze days (Liu and Diamond, 2005; Wang et al., 2013; Wang et al., 2014). Zhang et al. (2013)
52 suggest that chemical constituents and sources of PM_{2.5} in Beijing can largely vary with
53 seasons, which are characterized by variable meteorology and diverse air pollution sources.
54 For pollutions in Beijing, vehicles, coal combustion and cross-regional transfers are the
55 identical sources of PM_{2.5} (He et al., 2013). Nevertheless, specific meteorological conditions
56 play a key role in forming a haze weather phenomenon (Chen and Wang, 2015; Li and Zhang,
57 2016; Huang et al., 2014; Tang et al., 2015a). The severe haze event in January 2013 in
58 eastern China was closely related to the persistent weak surface winds and the atmospheric
59 boundary layer inversion near the surface (Zhang et al., 2014). By analyzing the haze episode
60 during 21-26 October 2014, Zhu et al. (2016) found that the severe air pollution in Beijing
61 was formed mainly by the southerly transport and strengthened by local contributions. The
62 main meteorological conditions during PHEs around Beijing include inversion in the
63 atmospheric boundary layer, weak winds near the surface and sufficient moisture in the air
64 (Liao et al., 2014). Wu et al. (2017) categorized two types of circulation conditions during
65 PHEs in Beijing–Tianjin–Hebei region: the zonal westerly type and the high-pressure ridge



66 type, giving rise to the descending air motion in the mid-lower troposphere, thus leading to a
67 lower boundary layer with higher concentration of pollutants. These studies have explored
68 ambient conditions in case studies. However, it remained unclear about the large-scale
69 atmospheric circulation backgrounds of PHEs around Beijing from a perspective of climate
70 change.

71 Recently, the role of underlying climatic factors in modulating regional weather conditions in
72 association with severe haze events has been reported (e.g., Niu et al., 2010; Wang et al., 2015;
73 Cai et al., 2017; Zou et al., 2017). These climatic factors, including the weakening East Asian
74 Winter Monsoon (EAWM) and related decreasing wind speeds (Niu et al., 2010), the
75 increasing relative humidity in the region (Chen and Wang, 2015), the reducing Arctic Sea ice
76 (Wang et al., 2015) and the anomalous SST in the subtropical western Pacific (Yin and Wang,
77 2016), etc., should have influenced the changes of severe hazes. The observed weakening
78 trend of the EAWM during the past few decades caused significant weakening winds over
79 North China, subdued the atmospheric transport of pollutants, and hence contributed to the
80 increasing number of haze days (Niu et al., 2010; Huang et al., 2012; Li et al., 2016). The
81 latest studies analyzed possible influence of anthropogenic climate change on haze
82 occurrences. Based on historical and future climate simulations, Cai et al. (2017) suggested
83 that the anthropogenic greenhouse gas emissions and the associated changes of the
84 atmospheric circulation might favor the haze weather conditions in Beijing. Zou et al. (2017)
85 indicated that the extremely poor ventilation conditions in eastern China could be linked to
86 the Arctic sea ice loss and the extensive Eurasia snowfall, which led to the unprecedented
87 severe haze event in January 2013. However, the linkage of underlying climatic factors to the
88 changes in PHEs around Beijing was still unclear, especially on long-term (multidecadal to
89 centennial) timescales. Specifically, it is interesting to explore how large-scale climatic
90 warming contributed to the increasing trend of PHEs in Beijing in the last decades.

91 In this paper, we analyze the characteristics of the PHEs in Beijing based on updated
92 observations, the associated atmospheric circulation changes related to the EAWM system
93 during 1980-2016. Furthermore, we explore the possible link between anomalous meridional
94 winds at 850 hPa in North China, as a practical index of EASM, and the SSTA over the
95 northwestern Pacific over the past centennial period 1900-2009. The data and methods used
96 are described in Sect. 2. In Sect. 3, we demonstrate the change of the PHEs in Beijing in the
97 past decades, the associated changes of meteorological conditions of the EAWM system and
98 investigate the linkage between anomalous meridional winds and SSTA in the northwestern
99 (NW) Pacific since 1900. Conclusions with more discussions are in Sect. 4.

100



101 2 Data and methods

102 2.1 Local visibility observations

103 The meteorological data used in this study are from the quality-controlled station observations
104 collected at the National Meteorological Information Center of China, including the relative
105 humidity, visibility, and weather phenomena. The data include 4 observations per day at 02:00,
106 08:00, 14:00, and 20:00 local time (LT). Consecutive records during the winters (December,
107 January and February, DJF) from 1980 to 2016 at 20 stations in Beijing are used.

108 In China, the visibility data at stations were obtained in different ways before and after 2013.
109 Before 23 January 2013, visibility was measured for meteorological purposes as a quantity
110 estimated by a human observer. Since then, the observations of visibility have been
111 transformed to instrumental measurements of the meteorological optical range (MOR). MOR
112 is defined as the length of the path in the atmosphere required to reduce the luminous flux of a
113 collimated beam from an incandescent lamp at a color temperature of 2700 K to 5 percent of
114 its original value; the luminous flux is evaluated by means of the photometric luminosity
115 function of the International Commission on Illumination (WMO, 2008). According to the
116 theoretical calculation (WMO, 1990; 2008), the transformation formula between the visual
117 estimate and the instrumental measurement is:

$$118 \frac{VIS_{\text{Instrumental}}}{VIS_{\text{Artificial}}} = \frac{(1/\sigma) \times \ln(1/0.05)}{(1/\sigma) \times \ln(1/0.02)} \approx 0.766 \quad (\sigma: \text{extinction coefficient})$$

119 It is necessary to address these data and maintain their consistency before analysis. In this
120 study, the visibility observations during 2013-2016 are transformed to be comparable with the
121 earlier visual estimations.

122 Haze is a multidisciplinary phenomenon, occurred occasionally. In different fields, haze is
123 represented by different variables, e.g. visibility and humidity in meteorology and PM_{2.5}
124 concentration in environmental science. In general, haze occurrences in meteorology are
125 defined based on the observations of relative humidity and visibility with specified criteria,
126 which vary among organizations (e.g. World Meteorological Organization and UK Met Office)
127 and empirical analyses (e.g., Wu, 2006; Vautard et al., 2009; Ding and Liu, 2014). In the
128 present study, we adopted a widely used comprehensive method based on weather
129 phenomenon, visibility and relative humidity. A haze day is defined if a haze weather
130 phenomenon is recorded with a daily mean visibility below 10 km and a daily mean relative
131 humidity below 90%. Early studies found that haze could be separated from fog by setting the
132 relative humidity to be below 90% (Schichtel et al., 2001; Doyle and Dorling, 2002). A
133 persistent haze event (PHE) is defined if haze is recorded at more-than-one sites in the region
134 for 4 consecutive days. The number of persistent haze days is calculated as the sum of the

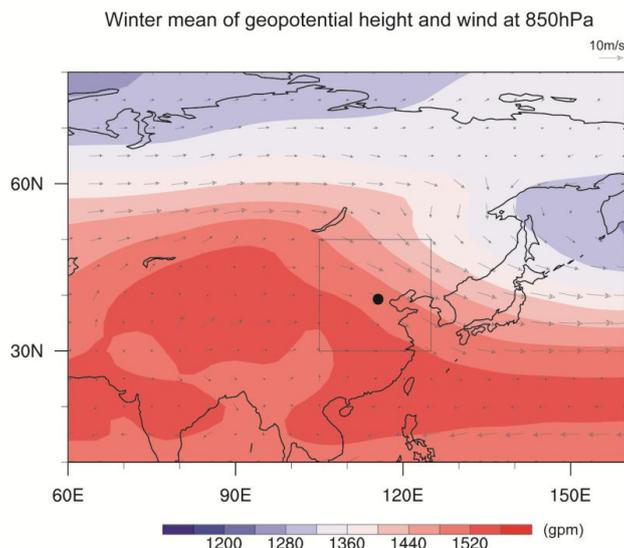


135 days during the PHEs.

136 **2.2 Global climate observations**

137 The daily and monthly data of the wind, geopotential height, specific humidity, sea level
138 pressure and air temperature from the NCEP/NCAR (National Centers for Environmental
139 Prediction/National Center for Atmospheric Research) for the period of 1980-2017, at 2.5 °
140 resolution (Kalnay et al., 1997) are used for the analysis of atmospheric conditions during the
141 PHEs. We also used the monthly data of meridional wind at 850 hPa for the period 1900-2010
142 from the 20th Century Reanalysis (20CR) version 2, at 2 °resolution (Compo et al., 2011) and
143 the ECMWF Atmospheric Reanalysis of the 20th Century (ERA-20C) at 1 °grid resolution
144 (Poli et al., 2016). The monthly SST data used are from the Hadley Center Sea Ice and Sea
145 Surface Temperature (HadISST) dataset, version 1.1, at 1 °resolution (Rayner et al., 2003) for
146 the period 1900-2011.

147 The dominant feature of the winter monsoon over East Asia is the northwesterly winds in the
148 lower troposphere (Fig.1). Severe haze usually occurs when the surface to mid-lower
149 troposphere northwesterlies weaken or even reverse to southerlies, indicating weakened
150 EAWM (Niu et al., 2010). The unusual southerlies at 850 hPa over eastern China were found
151 to be critical to the gathering of PM_{2.5} in Beijing (Cai et al., 2017). Thus, we proposed a
152 practical EAWM indicator, that is, the southerly winds at 850hPa over the region (30-50 °N,
153 105-125 °E) outlined in Figure 1. An extreme southerly day is defined if the regional mean
154 anomalous 850 hPa meridional winds exceeds 2 standard deviations above the boreal winter
155 average in the region.



156

157 **Fig. 1. Climatological mean of winter wind (vector, units: ms^{-1}) and geopotential height**
158 **(shading, units: gpm) at 850 hPa over East Asia during 1980-2016. The outlined region**
159 **(30-50° N, 105-125° E) is used to calculate a regional mean climatological background**
160 **around Beijing (black dot).**

161 2.3 Ensemble Empirical Mode Decomposition (EEMD)

162 In present paper, the EEMD method is applied to separate the multidecadal-to-centennial
163 timescale variations of the SSTA time series over the northwestern Pacific. The EEMD
164 method is an adaptive time-frequency data analysis technique developed by Wu et al. (2007)
165 and Wu and Huang (2009). It is an efficient way to separate specific timescale variations in
166 the original data series. The EEMD method is a refinement of the Empirical Mode
167 Decomposition (EMD) method, which emphasizes the adaptiveness and temporal locality of
168 the data decomposition (Huang et al. 1998; Huang and Wu 2008). With the EMD method, any
169 complicated data series can be decomposed into a few amplitude-frequency modulated
170 oscillatory components, called intrinsic mode functions (IMFs), at distinct timescales.

171 The main steps of the EEMD analysis in this study are as follows (Qian, 2009): (1) add a
172 random white noise series with an amplitude 0.2 times the standard deviation of the data to
173 the target time series to provide a relatively uniform and high-frequency extreme distribution,
174 allowing the EMD to avoid the effect of potential intermittent noise in the original data; (2)
175 decompose the data with the added white noise into IMFs using EMD; (3) repeat steps 1-2 for

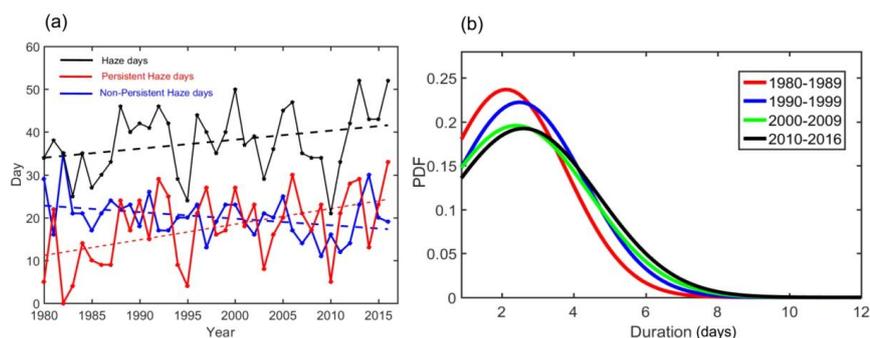


176 1000 times, but with distinct random white noise series added each time; and (4) obtain the
177 mean IMFs of the 1000 ensemble results to produce the final results.

178 Via EEMD, the mean SSTA series over the northwestern Pacific (120°E – 150°E , 26°N – 40°N)
179 for the winter during 1900–2009 is decomposed into a nonlinear secular trend (ST) and 5
180 major timescales of IMFs (figure not shown). The multidecadal variability (MDV) is
181 represented by the fifth IMF, with an oscillatory period between half and one century.

182 3 Results

183 3.1 Increasing PHEs in Beijing from 1980 to 2016

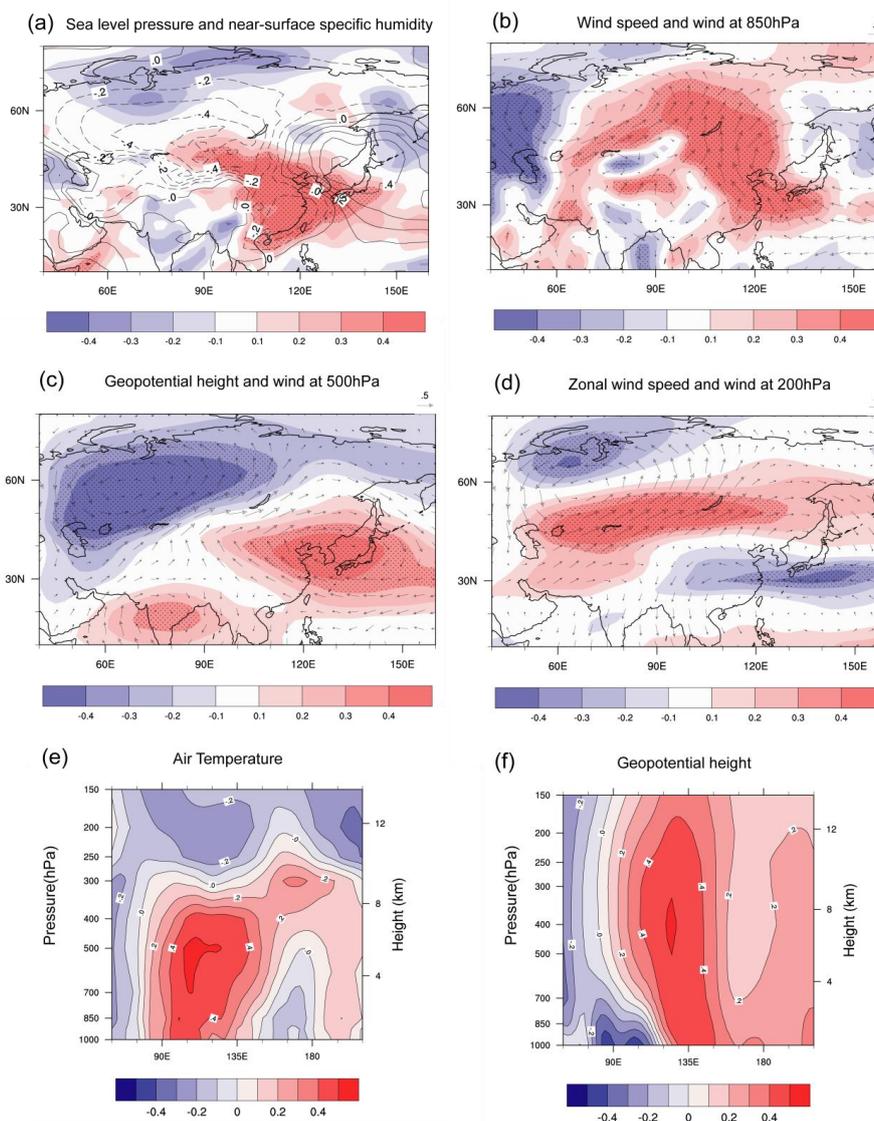


184

185 **Fig. 2. (a) Number of haze days (black), persistent haze days (red) and non-persistent**
186 **haze days (blue) in winter in Beijing during 1980–2016. Dashed lines show the least**
187 **squares trends. (b) The Probability Distribution Function (PDF) of the duration (unit:**
188 **days) of haze events in Beijing for each decadal period during 1980–2016.**

189 As shown in Fig. 2a, the number of haze days in Beijing exhibits large inter-annual variability
190 with a non-significant increasing trend (black curve in Fig. 2a), consistent with previous
191 studies (e.g., Chen and Wang, 2015). However, the number of persistent haze days in Beijing
192 exhibits a significant increasing trend, while that of the non-persistent haze days exhibits a
193 significant decreasing trend, significant at the 0.05 level. Therefore, it is the number of
194 persistent haze days that has been increasing during the past decades. The PDF of the duration
195 of haze events in Beijing (Fig. 2b) indicates that most haze events lasted for about 3 days.
196 However, the probability of haze events lasting for 1–3 days decreased, while that of the
197 longer events increased from the 1980s to the present. The PHEs lasting for more than 4 days
198 have increased remarkably. The duration of haze events have tended to be longer in the past
199 decades.

200 3.2 Meteorological conditions for PHEs



201

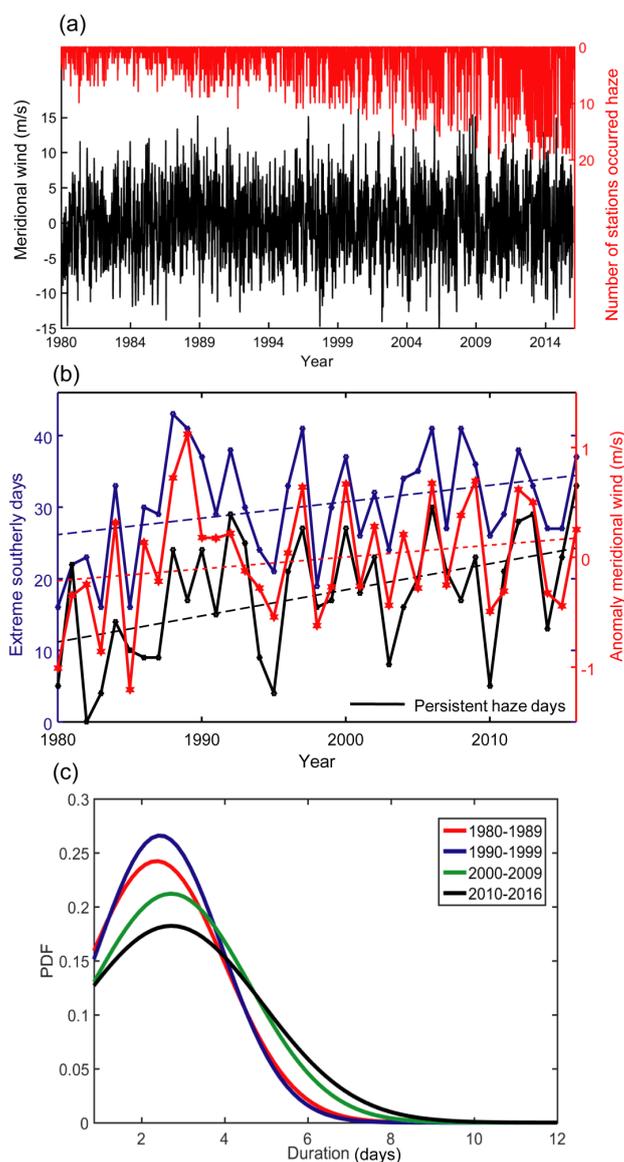
202 **Fig. 3.** The correlation coefficients between the number of persistent haze days and (a)
203 sea level pressure (contour) and near-surface specific humidity at 1000 hPa (shading); (b)
204 wind (arrows) and wind speed (shading) at 850 hPa; (c) wind (arrows) and geopotential
205 height (shading) at 500 hPa; (d) wind (arrows) and zonal wind speed (shading) at 200
206 hPa; (e) vertical profile of air temperature at 40 °N; and (f) vertical profile of
207 geopotential height at 40 °N in winter during 1980-2016, based on the monthly
208 NCEP/NCAR reanalysis data. The linear trend is removed before calculating the



209 **correlation. The black dots indicate significant changes at the 0.05 level using the t-test.**

210 Figure 3 depicts the correlation coefficients of the anomalous variables of atmospheric
211 circulation from near-surface to upper troposphere with the number of persistent haze days in
212 Beijing in the winter during 1980-2016. In the lower troposphere (Fig. 3a), most of China is
213 covered by an anomalous low-pressure system, adjacent to an anomalous high over the
214 western Pacific, suggesting weaker-than-usual northerly winds from the mid-high latitudes.
215 Consequently, North China is covered by the anomalous southerlies, resulting in a significant
216 decrease in wind speed (Fig. 3b). Southerly winds bring warm and moist air from the south to
217 the north part of China, creating favorable moisture conditions for haze occurrences (Fig. 3a).
218 At 500 hPa, East Asia is mainly dominated by an anomalous high (Fig. 3c), representing a
219 shallow East Asian Trough. The associated northwesterlies exist to the north rather than the
220 south of Beijing, limiting the cold and dry northwesterly flows to Beijing and reducing the
221 wind speed in Beijing. These atmospheric circulation patterns are favorable for severe haze
222 occurrences, as partly discussed in previous studies (Chen and Wang, 2015; Yin and Wang,
223 2016; Cai et al., 2017). At the upper troposphere 200 hPa, the East Asian jet stream shifts
224 northward (Fig. 3d) with enhanced zonal circulation in the high latitudes and weakened
225 meridional circulation over East Asia. This pattern indicates weak cold air activity around
226 Beijing. The decreased zonal wind in the middle latitudes favors the maintenance and
227 enhancement of the pollutant convergence needed for the haze occurrences. The vertical
228 profiles of air temperature and geopotential height at 40°N around Beijing suggest an
229 abnormal pressure system over the northwestern Pacific region (40°N, 120-155°E) with
230 anomalously warm temperature and high geopotential height from 850 hPa to 300 hPa (Fig.
231 3e and f). This anomalous vertical structure of atmosphere is also favorable for stagnant
232 weather conditions and gathering of pollutants in the atmospheric boundary layer around
233 Beijing.

234 **3.3 Variations of extreme southerlies and its relationship with PHEs**



235

236 **Fig. 4.** (a) Number of stations where haze is recorded in Beijing (red) versus daily
237 normalized meridional wind at 850 hPa over the study region (30-50 °N, 105-125 °E)
238 (black) in winter during 1980-2016. (b) Anomalous meridional wind (red), number of
239 extreme southerly days (blue) and persistent haze days (black) in winter during
240 1980-2016. Dashed lines show the least squares trends. (c) The Probability Distribution
241 Function (PDF) of the duration (unit: days) of extreme southerly episodes for each

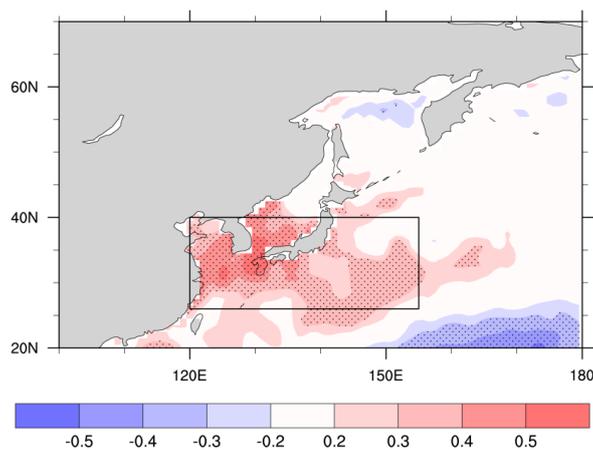


242 **decade during 1980-2016.**

243 The weakening of the EAWM system was considered to be responsible for the changes of
244 meteorological conditions in Fig. 3 (Niu et al., 2010; He, 2013; Wang and He, 2013; Yin and
245 Wang, 2016). During PHEs, the northerly winter monsoon weakened and brought less cold
246 and dry air to the region, favorable for the formation and maintenance of haze. According to
247 present analysis, the meridional wind anomalies at 850hPa in North China could be one of the
248 most effective conditions for haze occurrence.

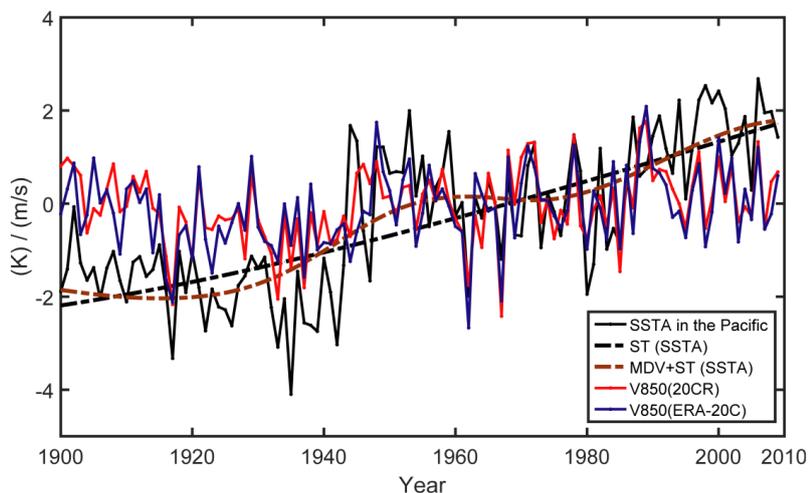
249 As shown in Fig. 4a, daily meridional wind anomalies are notably correlated with the number
250 of haze stations in winter during 1980-2016. The correlation coefficient is 0.43, at $\alpha = 0.01$
251 level. It suggests that the southerlies well correspond to the haze occurrence in Beijing. The
252 time series of the meridional wind anomalies in winter exhibits strong interannual variability
253 with a non-significant increasing trend (red lines in Fig. 4b). However, the number of extreme
254 southerly days exhibits a significant increasing trend (significant at the 0.05 level) and it has a
255 significant positive correlation coefficient of 0.70 (0.66) with the number of persistent haze
256 days in Beijing with (without) trends, significant at $\alpha = 0.01$ level. The PDF of the duration of
257 extreme southerly events indicates that most of the extreme southerly episodes lasted for 3
258 days (Fig. 4c). However, the probability of the extreme southerly episodes lasting for 1-3 days
259 is decreasing, while that of the longer episodes is increasing. The change in the extreme
260 southerlies corresponds well to that of the persistent hazes in Beijing. Clearly, the higher
261 probability of the longer extreme southerly episodes was underlying the increasing
262 occurrences of PHEs in Beijing.

263 **3.4 Linkages between SSTA over the NW Pacific and meridional wind anomalies since**
264 **1900**



265

266 **Fig. 5.** The correlation coefficients between SSTA in the northwestern Pacific and the
 267 number of extreme southerly days in winter in North China during 1980-2011. The
 268 linear trend is removed before calculating the correlation. The black dots indicate
 269 significant correlation at the 90% confidence level using the t-test.



270

271 **Fig. 6.** Time series of the normalized SSTA in the northwestern Pacific (black),
 272 meridional wind anomalies at 850 hPa in North China from 20CR (red) and ERA-20C

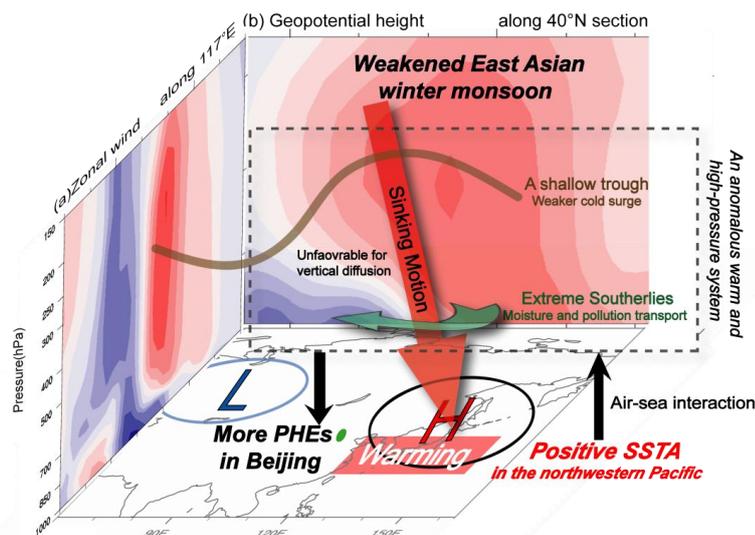


273 **(blue) during 1900-2009. The climatic mean is calculated for the period 1961-1990. The**
274 **black dotted curve is the secular trend (ST); the brown dotted curve is the combination**
275 **of ST and the multidecadal variability obtained via EEMD.**

276 As shown in Fig. 5, there is a positive SSTA zone in the subtropical northwestern Pacific
277 (120–155 °E, 26–40 °N). This suggests that the northerly winter monsoons in East Asia
278 become weaker, with more extreme southerly episodes, when the subtropical northwestern
279 Pacific is warmer. It is interesting to investigate this relationship over a longer period. As
280 shown in Fig.6, over the past centennial period 1900-2009, SSTA in the subtropical
281 northwestern Pacific and the meridional wind anomalies at 850 hPa over North China well
282 co-varied, especially at the multidecadal timescale. The correlation coefficients are 0.46
283 (detrended: 0.42) and 0.51 (detrended: 0.53), based on the ERA-20C and 20CR data,
284 respectively, all significant at $\alpha = 0.01$. During this period, the SSTA showed a warming
285 secular trend. The combination of MDV and ST of SSTA exhibits a particular warming phase
286 since the mid-1980s. As discussed above, this notable warming phase in the subtropical
287 Pacific could lead to a weakened EAWM, with increasing number of extreme southerly
288 episodes, and hence increasing PHEs in Beijing.

289 **4. Conclusion and discussion**

290



291

292 **Figure 7. Schematic diagrams summarizing the dynamical linkage between positive**
293 **SSTA in the northwestern Pacific, via the weakening East Asian winter monsoon system,**
294 **and increasing PHEs in Beijing. The cross sections are correlation coefficients between**
295 **meridional wind anomalies and vertical profile of (a) zonal wind at 117 °E and (b)**
296 **geopotential height at 40 °N at 500 hPa in winter during 1980-2016. The latitude and**
297 **longitude coordinates of Beijing (green dot in base map) is 40 °N, 117 °E. The letter “L”**
298 **and “H” in base map demonstrate anomalous low-pressure and high-pressure system,**
299 **respectively.**

300 So far we have demonstrated the increasing frequency of PHEs in Beijing based on updated
301 observations and the associated changes in large-scale atmospheric circulations during
302 1980-2016, and then explored the linkages of sea surface temperature in the northwestern
303 Pacific in the past centennial period 1900-2009. The essential meteorological conditions
304 favorable for the increasing PHEs in Beijing are associated with the weakening East Asian
305 Winter Monsoon system, which is partly a consequence of the warming (positive SSTA) over
306 the NW Pacific via air-sea interaction. Figure 7 illustrates such a dynamic mechanism.

307 There is a secular warming trend over the subtropical northwestern Pacific, consequently with
308 anomalous high pressure or anti-cyclone system in the mid-lower troposphere over the region
309 maintained via air-sea interaction. It would induce anomalous southerlies over the East Asian
310 continent. Therefore, as a consequence, the EAWM system weakens. Particularly, in the lower
311 troposphere, the weakening monsoons are more likely than before interrupted by persistent
312 southerlies in North China, which bring warm and moist air from the south to the north of



313 eastern China and create favorable moisture conditions for haze occurrences. In the
314 mid-troposphere, the East Asian Trough tends to be shallower and prevent the associated
315 northwesterlies from expanding towards the south of Beijing, hence favorable for gathering of
316 pollutants around Beijing. These anomalous circulation patterns tend to cause sinking air
317 motion in the mid-lower troposphere over Beijing, leading to stagnant weather conditions and
318 gathering of pollutants in the atmospheric boundary layer. These changes in SSTAs and
319 associated atmospheric circulation clearly contribute to the increasing PHEs in Beijing.

320 One of the most direct circulation factors for haze occurrences is weakened northerlies in the
321 lower troposphere around Beijing. During 1980-2016, the change in the extreme southerlies
322 corresponds significantly to that of the persistent hazes in Beijing. The higher probability of
323 the longer extreme southerly episodes in the past decades was underlying the increasing
324 occurrences of PHEs in Beijing. Furthermore, we found that over the past centennial period
325 1900-2009, the anomalous meridional winds at 850 hPa were positively correlated with the
326 SSTA over the subtropical northwestern Pacific. Since the mid-1980s, a notable warming
327 phase in the northwestern Pacific has weakened the East Asia winter monsoon system,
328 increased the number of extreme southerly episodes in eastern China, and therefore
329 contributed to the increasing PHEs in Beijing.

330 During the past centennial period 1900-2009, the SSTA in the northwestern Pacific showed a
331 notable warming secular trend. In fact, the northwestern Pacific is one of the most stable
332 warming regions, as a part of global warming during the last century (Zeng et al., 2001).
333 IPCC (2013) has concluded that it is very likely that anthropogenic forcings have made a
334 substantial contribution to the increase of global upper ocean heat content (0–700 m) as
335 observed since the 1970s. Based on the results of 15 models from the Coupled Model
336 Intercomparison Project Phase 5 (CMIP5), Cai et al. (2017) projected some circulation
337 changes induced by increasing atmospheric greenhouse gases that might contribute to the
338 increasing hazes in Beijing. The present paper suggests a more concrete observation-based
339 mechanism for explaining how the changes in atmospheric circulation contribute to the
340 increasing haze occurrences in Beijing.

341 Nevertheless, emissions of pollutants should also have contributed to increasing hazes in
342 Beijing. It remains interesting to quantify contributions of pollutants emissions and climate
343 change to the occurrences of PHEs in Beijing. Further studies are needed.

344

345 **Data availability.**

346 Atmospheric circulation data are available from the NCEP/NCAR data archive



347 (<http://www.esrl.noaa.gov/psd/data/gridded/data.ncep.reanalysis.html>) and ECMWF data
348 archive (<https://www.ecmwf.int/en/forecasts/datasets/browse-reanalysis-datasets>). SST data
349 are downloaded from Met office Hadley centre observation datasets
350 (<https://www.metoffice.gov.uk/hadobs/hadist/>). The ground observations are from the
351 National Meteorological Information Center of China (<http://data.cma.cn/>). The atmospheric
352 composition data can be obtained from the authors.

353

354 **Competing interests:** The authors declare that they have no conflict of interest.

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