

We would like to express our sincere thanks to the two reviewers for their careful reading and constructive suggestions, which have helped us improve the quality of this manuscript. We have addressed all their comments carefully and revised the manuscript accordingly. The detailed responses to their concerns and comments are presented as follows.

Anonymous Referee #1

This paper summarizes the overall performance of several fully-coupled air quality models that participated in the MICS-Asia III intermodal comparison study. It is the first part of a multi-part study. While the paper is well organized and the discussion is straight-forward, there are numerous places where the grammar needs to be fixed. I have tried to make some suggestions in my specific comments; however, the authors should not assume I have found all the problems. There are some aspects of the manuscript that are not explained well, such as the rationale for this paper in relation to future parts and the rationale of the intercomparison framework. By the time I finished the paper, I feel like I did not learn anything substantially new; therefore, the authors have not adequately highlighted the new results of this study.

Major Comments: 1) In the introduction, the authors talk about Topic 3 of MICS-Asia phase 3 which is the subject of this paper. At the end of the paper I felt like I did not get any information regarding the feedbacks. Perhaps the paper title is implying that those details will be included in subsequent parts. It would be useful at the end of the introduction to have a clear understanding of what the objectives are for this paper, versus subsequent parts that will appear.

Response: Thanks for this great suggestion. In the last paragraph of introduction, we claim that “This paper presents an overview of the MICS-ASIA III Topic 3, serving as the main repository of the information linked to Topic 3 simulations and comparisons”. To make it clear, we added in the manuscript: “Specifically, this paper aims to archive the information of participating models, how the experiments, and results of model evaluation. The results of the MICS-Asia Topic 3 experiments looking at the direct and indirect 135 effects during heavy haze events will be published in a companion paper, part II.”

2) The authors do speculate why there are differences among the models; however, the paper would be improved significantly if they went into more detail into a few instances to find more concrete reasons for the differences. This might require more analysis of the results. But as the

paper stands, it does not shed any new light on why the air quality models could differ. In the conclusion, the authors state that the paper provides “some directions of future model developments”, but I see no evidence of that in the paper. The authors could also do a better job at citing papers that examine processes that might be missing or poorly represented in the air quality models.

Response: Thank you for pointing out. We added more explanations on why model differs in the revised manuscript. For example, the radiation differences in M6 is due to the use of different meteorological boundary conditions, and the averaged values were taken from 3hourly outputs. More importantly, high liquid water path in M6 lead to lower radiation from M6 in north China.

We have cited more papers on the reasons for underestimation of sulfate and SOA, and clearly claim the future directions for future model developments: “X. Huang et al. (2014) found including natural and anthropogenic mineral aerosols can enhance sulfate production through aqueous-phase oxidation of dissolved SO₂ by O₃, NO₂, H₂O₂ and transition metal. Gao et al. (2016), Wang et al. (2014), and Zhang et al. (2015) also emphasized the importance of multiphase oxidation in winter sulfate production. However, these processes are currently not incorporated in the participating models for this study, which might be responsible for the apparent under-predictions of sulfate concentration”;

“An et al. (2013) incorporated photoexcited nitrogen dioxide molecules, heterogeneous reactions on aerosol surfaces, and direct nitrous acid (HONO) emissions into the WRF-Chem model and found these additional HONO sources can improve simulations of HONO and nitrate in north China. M7 also predict high nitrate concentrations (N₂O₅ and NO₂ gases react with liquid water, Zheng et al., 2015), and the predicted lower nitrate concentrations from other models are probably due to missing aqueous phase and heterogeneous chemistry, or the implementations of different gas phase oxidation in these models. Many studies have been conducted regarding sulfate formation issues. Nitrate also account for a large mass fraction in PM_{2.5} during winter haze events in north China, yet less attention was attracted to fully understand its formation. It is worth furtherly digging into the details about how different processes contribute to high nitrate concentrations in future studies”;

“R. Huang et al. (2014) also suggested that low temperature does not significantly reduce SOA

formation rates of biomass burning emissions. Most models over-simplified SOA formation”; “Zhao et al. (2015) comprehensively assessed the effect of organic aerosol aging and intermediate-volatile emissions on OA formation and confirmed their significant roles. All these results suggest more complicated SOA scheme are needed to improve organic aerosol simulations during haze events”.

3) The purpose of the MICS-Asia phase 3 is to look at feedback effects. I gather that is not that subject of this paper, and this paper is showing the initial evaluation of the aerosol simulations that will be important when looking at feedback effects later. The authors go into some detail on evaluating aerosol composition, but do not say anything about size distribution. Size distribution will be just as important for optical and CCN properties. I suggest adding a section that compares the predicted size distribution in some manner. I assume there are some sort of size distribution measurements that could be utilized. If not, it would still be useful in the context of the subsequent papers.

Response: Thanks for this good suggestion. We are aware that size distribution is crucial to optical properties and CCN formation, and thus direct and indirect effects. The participating models use mode and sectional approaches with different degree of complexity, we present the basic information on size distribution in each model, such as geometric mean radius and standard deviation and the number and range of size bin, and briefly mention about its potential influence on AOD simulation. We know this is not enough and we will investigate this issue in more detail in the companion paper (part II) by conducting additional sensitivity simulations with different size distribution while keeping other conditions the same. For this paper, we add more information on aerosol size distributions in the models in Table 1. Except M2 using 8-bin MOSAIC sectional approach, the other models use mode and bulk approach, with similar mean radius and standard deviation for anthropogenic aerosols or no size information. Unfortunately, we don't measurement of size distribution for comparison.

Specific Comments:

Line 43: Change “resolutions” to “resolution”.

Response: We have changed.

Line 56: I would change “are consistent” to “are similar”. “consistent” can imply that the model results are good, but they could all consistently disagree with data. Change “haze event” to “haze events”.

Response: We have changed.

Line 58: “some brief senses” is an awkward phrase and should be replaced. The abstract could be shortened so that it contains only the most important findings. For example, the sentences in lines 44 – 48 could be removed. The whole abstract is rather weak.

Response: We have changed “provide some brief sense of” to “present”; we remove lines 44-48. This abstract is mainly used to introduce Topic 3 and how simulations were done and analyzed. Major findings out of Topic 3 will be published in the following companion papers.

Line 66: Change “but primarily in Asia” to “but most deaths occur primarily in Asia.”

Response: We have made this change.

Line 74: I would not use a semi-colon here and just have two sentences, although the second would need to be rephrased slightly.

Response: We have made this change.

Lines 134-139. There appears to be no underlying motivation for how the air quality models are compared. The only constraint on the models was the use of the same emissions inventories and they had to provide a set of variables. To better isolate the differences among the models, it would have been useful to have similar domains, grid spacings, and boundary conditions. I understand it would make the setting up the models a bit more difficult, but it would significantly reduce the differences arising from boundary conditions and spatial resolution. There are already many differences associated with the internal treatments of meteorology, chemistry, and aerosols. What I am looking for here is some further explanation as to why MICS-Asia organizers found the current configuration sufficient.

Response: We tried our best to constrain the differences, by requesting the modeling groups to use the same emission inventories, domains, grid size and boundary conditions, but some models has been constructed and tested on their own configurations, such as M5 and M6, and all the modeling groups use a similar resolution ~50km (Table 1) except M4 (15km). M4 actually covers a nested domain of M3. We include this simulation to further look at the impacts of grid spacing on aerosol feedbacks, which will be documented in the companion paper.

Models are quite different from each other, and it is difficult to keep all the inputs the same. Even using the same boundary conditions, mapping species from global chemistry simulations to different chemical mechanism each model uses can also lead to differences. We describe what we provided in the methodology section was intended to reduce the differences in inputs. Results from these model configurations have been published to report aerosol feedbacks and aerosol direct and indirect effects in Asia. The underlying motivation is to look at how current reported values differ, and provide possible range of these values. Given this motivation and our efforts on constraints, we think our current configuration could be sufficient although not ideal.

Line 153: The Grell reference is correct, but it only describes the initial model which did not discuss any of the feedback processes – which seems to be the focus of MICS-Asia Phase III. Those feedbacks were first implemented in Fast et al. (2006) and revised in subsequent manuscripts.

Response: We have added Fast et al. (2006).

Line 182: For VBS, need to cite Adhamov et al. (2012) in which it was developed and described.

Response: We confirmed that M1 did not use VBS, so we deleted description here.

Lines 190-191: The sentence regarding SOA is not correct and misleading. A VBS SOA treatment has been available in the public version of WRF-Chem for several years. What the authors mean to say is that the version of MOSAIC used in this study includes no SOA. The correct language here should imply that the users have chosen not to include SOA.

Response: We changed it to “The MOSAIC version used in M2”.

Section 2.1, in general. The description of the models is uneven. Some sections go into some details about the aerosol model, such as noting the geometric means of the modes (e.g. M5) but not going into the same detail as another model (e.g. M1). For one model the details of how optical properties and hygroscopicity are discussed, but then another model the same level of detail is not discussed. The authors need to revise this section to have the appropriate level of detail for all models.

Response: Thanks for these good suggestion. We reorganize the whole section to make it easier to read and keep details for all models even.

Lines 232 – 239: The text discusses differences in the physics configurations as it should, but I assume the other models have physics differences too. Why not state that?

Response: Thanks for these good suggestion. We have added other physical configurations in Table 1 and in the texts.

Line 286: What does regriding mean? To handle the emissions inventories properly, the emissions need to be reapportioned so that mass is neither gained or lost. Regriding implies interpolation, that to me indicates a poor method of handling the emissions from one domain to another.

Response: Yes, the emissions were handled using mass conservative method to make sure that mass is neither gained nor lost, not simple interpolation.

Line 409: Change “are frequently happening” to “frequently happen”.

Response: We have changed.

Section 4, in general: this study relies on comparing model output to relatively few (at least for the PM data) point measurements. However, some discussion is needed to put the proper context of this type of comparison since the grid size differs among the models so there are issues of representativeness that must be considered.

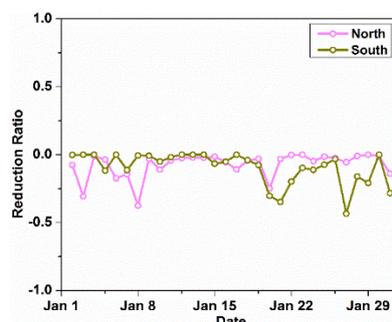
Response: This is a good point. We required the modeling groups use similar model resolutions. For most models, ~50km resolution is used. Except model evaluation, the model resolution would also affect simulations of meteorological fields and then chemical fields. We are preparing a manuscript discussing the influences of model resolution by keeping all other factors the same, which will be published in this topic series. Based on this, we think the comparison representativeness would be close for participating models. For model observation comparison, this is a common problem, given limited available observations in this study. In our previous experiences, the comparisons are done mostly in urban regions, and measurements during haze events show high homogeneity when we compared concentrations from different city sites in previous study. Besides, the in situ measurements at city sites exhibit a similar variation trend during haze events, reflecting haze pollution at regional scale rather than local scale, thus the concentration difference resulting from grid size would reduce.

Line 442: Turbulent mixing is missing from this description, which is not the same as transport.

Response: We have added turbulent mixing.

Lines 469-483: What is missing from this discussion is how clouds affect the prediction of downward shortwave radiation. I assume that the clouds are the main factors controlling clouds, but there is no mention of this. Would be useful to include what the clear-sky values are in Figure 5.

Response: Thanks for this great point. Since clear sky radiation is not submitted by modeling groups and it is not a default model outputs, it might not be appropriate to show it in Figure 5. However, we do think it is important to include the discussion here because we agree with you that it is one of the main factors controlling radiation. We add a figure to show the reduction ratio of downwards shortwave radiation due to clouds $(SW-SW_{clearsky})/ SW_{clearsky}$ in the SI, and discussion on clouds: “Clouds are also important to alter radiation. To exclude its impacts on the radiation shown here, we calculated the reduction ratio of radiation due to clouds. During the severe haze period (16-19 January 2010), the averaged reduction fraction is 5.9% in north China and 4.2% in south China, suggesting the relatively lower radiation during this period shown in Figure 5(d) is mainly caused by aerosols, while the lowest radiation on 20 January was caused by clouds (Figure 5(d)).”



Line 493: Change “larger near surface” to “larger near the surface”.

Response: We have changed.

Line 498: Awkward sentence – need to revise.

Response: We removed “only”.

Line 513-514: Change to “All models produce similar CO predictions” based on how I understand this sentence.

Response: We have changed.

Line 522: It is rather surprising that the models produce better PM than ozone. Usually it is the other way around.

Response: This is probably because previous good ozone simulations mostly occur in summer. The study period is frequently affected by heavy haze in winter, when photochemistry is weaker and other chemical processes maybe more complicated. Another reason is that primary PM (including POA, BC etc) is of similar magnitude to secondary aerosol in winter haze period, which means meteorological and chemical processes are equally important.

Line 548: I doubt that sea-salt emissions are responsible for differences in PM10.

Response: We recheck the model and possible reason for the differences in PM10. Fig 7d shows that M5 largely overpredict PM10 concentration during 1-8 January at Rishiri, where is less affected by continental aerosols, overestimation of sea salt emission under certain meteorological conditions (such as large wind) is responsible for such positive biases, whereas in the rest of January, M5 perform relatively better and is generally consistent with most of the models.

Line 569-580: Chen et al., ACP (2016) is just but one paper that describes possible missing reactions associated with sulfate. It would be useful for the authors to delve a bit deeper into the literature to find such issues associated with models. Most community models are inherently dated and do not necessarily have the most up-to-date chemistry treatments since it takes time for new research findings to make there was in to those community models.

Response: We have added several papers describing the model problems and how missing mechanism can explain the underestimation of sulfate and other model problems. All the participating models have not implemented such mechanism, which is an important source for the underprediction of sulfate in this study.

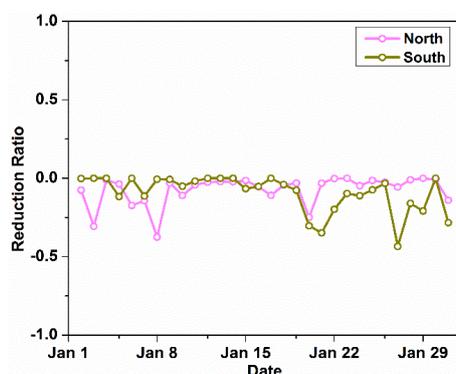
Line 593-594: There are probably other reasons as well for errors in nitrate predictions.

Response: Thanks for mentioning this. We added the spatial plots of NO_x in the SI. It is shown that NO_x from M5 is the highest, yet the nitrate produced from M5 is lower than other models except M1, which suggest there could be some missing nitrate formation pathways or stronger deposition of nitrate and its precursors in M5. Other than the mentioned aqueous phase and heterogeneous chemistry, the implementation of gas phase chemistry in different models might also played a role here. We added this to the revised sentence: “, or the implementations of different gas phase oxidation in these models. Many studies were conducted for sulfate formation issues. Nitrate also account for a large mass fraction in PM_{2.5} during winter haze

events in north China, yet less attention was attracted to fully understand its formation. It is worth furtherly digging into the details about how different processes contribute to high nitrate concentrations in future studies.”

Line 607: The authors list deposition, but this usually means dry deposition. What about wet scavenging? Same comment applies to line 610.

Response: Thanks for this point. Here we meant both dry and wet deposition. To make it clearer, we added “(dry deposition and wet scavenging)”. We added a figure in the SI to check how often were clouds present during this month. There were clouds only in a few days, but they can be very important for wet scavenging for hydrophilic aerosols. Thus, we added wet scavenging in Line 607 and 610.



Line 607: Change predicted BC to “predicted BC at the surface”

Response: We have changed.

Line 609: I think the authors mean horizontal grid resolution and not “horizontal grid interpolation.” I have no idea what the latter means in this context. Please be more specific.

Response: We have changed.

Line 612: Since POC is about the same from the models, then BC should be as well. So it is a bit of a mystery why BC from M2 and M7 are higher than the other models.

Response: It depends on the models treat deposition and aging processes. For example, in the GOCART aerosol model (M3 and M4), 80% of BC are assumed to be hydrophobic and then undergo aging to become hydrophilic in an e-folding time of 1.2 days. Hydrophilic aerosols will go through wet deposition. But in other models like M2 and M7, BC is assumed to be hydrophobic, thus the wet removal is less activated. We added these explanations in the revised manuscript.

Line 621: Find a reference for this comment – there are lots of papers to cite here.

Response: Thanks. We added the following citation.

Heald, C. L., Henze, D. K., Horowitz, L. W., Feddema, J., Lamarque, J.-F., Guenther, A., Hess, P. G., Vitt, F., Seinfeld, J. H., Goldstein, A. H., and Fung, I., 2008. Predicted change in global secondary organic aerosol concentrations in response to future climate, emissions, and land use change, *J. Geophys. Res.*, 113, D05211, doi:10.1029/2007JD009092.

Line 638: This implies the model is missing a feedback, and I thought this study was about in the inclusion of feedbacks (see line 106 on topic 3).

Response: Yes, we are trying to say less VOCs should be able to convert to SOA under hazy conditions, but the model still used 10% yield, which could overestimates SOA.

Line 642: “dust deflation” is an odd phrase. What is that?

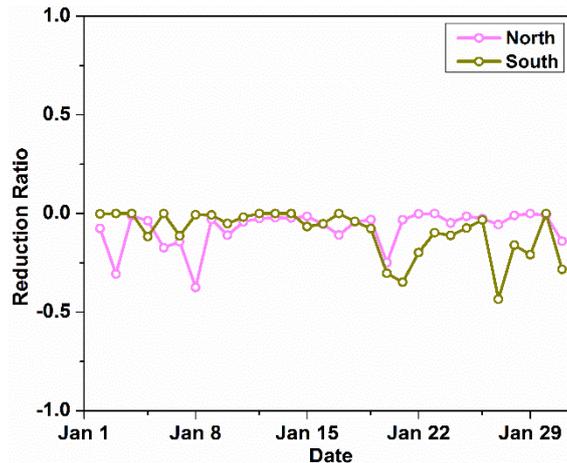
Response: We change to “wind-blown dust”

Lines 650-652: These sentences are poorly written. Suggest changing to “Only the sulfate predictions from M5 are close to the observed values. Sulfate is much lower than observed for all other models, except M6 which is too high. M2 and M7 predict reasonable nitrate concentrations. M3 and M4 overpredict OC during the haze period, but other models underpredict OC concentrations.”

Response: We have changed following your suggestions.

Line 677: What about clouds? How often were AOD retrievals not possible due to cloudy conditions?

Response: Yes, it is possible. But in winter, there is relatively less cloud amount, so cloud is not a serious problem to AOD retrieval. For example, in Figure 14 and SI, about three days within a month (Jan. 2 Jan. 8 Jan 20) are due to clouds. We have added the influences of clouds in the text: “under serious pollution and cloudy conditions”. Thanks for pointing out.



Line 680: The figure captions should also state that the AOD is a daily (daytime) value.

Response: We have changed following your suggestions: “daily (daytime) mean”

Line 686: Change “it’s” to “it is”.

Response: We have changed following your suggestions.

Line 718: Change “shows overprediction” to “shows an overprediction”.

Line 724: Change “lower RH simulation“ to “lower simulated RH”.

Response: We have changed following your suggestions.

Line 726: Change “OC concentration” to “OC concentrations”.

Response: We have changed following your suggestions.

Line 736-737: The authors have not shown this. It is very likely the size distribution and mixing state is treated differently. In this sense, the explanation provided previously in the paragraph is incomplete. I doubt one can really attribute the difference in AOD without a more rigorous analysis than the simple explanations presented here. At best, they are showing the range of AOD associated with all the differences among the models.

Response: Thank you for this question and suggestion. We added the mixing state of each model in Table 1. Only M6 used external mixing. Curci et al. (2015) discussed the impacts of mixing state on simulated AOD and found that external mixing state assumption significantly increase simulated AOD. M6 used external mixing but shows a relative lower AOD because of ignorance of other aerosol species like dust, sea-salt, etc.. Other models used internal mixing for major aerosol compositions. The size distribution treated in the models except M2 (sectional approach) and GOCART bulk approach is described by a lognormal distribution with similar geometric mean radius and standard deviation for different modes or species, such as

0.07 μm for inorganic aerosols, 0.01 μm and 0.02 μm for BC and OC, but different bins for dust and sea salt (their concentrations are low in the north China Plain in winter). Thus we believe the differences in AOD shown here is mostly due to differences in simulated compositions. The size distribution in this sentence means the used modal size treatments in M1, M5, M6 and M7. To avoid misunderstanding, we change it to “lognormal treatments” and add the conclusion from Curci et al., (2015) about the impacts of mixing state.

Curci, G., et al. "Uncertainties of simulated aerosol optical properties induced by assumptions on aerosol physical and chemical properties: An AQMEII-2 perspective." *Atmospheric Environment* 115 (2015): 541-552.

Lines 773-775: I don't see how interpolation of emissions to the grid should lead to model uncertainties. Of course, there could be errors introduced to reapportion emissions from one grid to another. But these would only be large if the mathematical method of reapportionment is poorly treated. There are ways to ensure that such uncertainties are small.

Response: Thanks. We changed this sentence to “which might be caused by the treatment of aging and deposition (dry deposition and wet scavenging) processes.”

Line 776: “Manifold” is a strange word to use in this context.

Response: We have changed to “various”.

Line 783: And what are those improvements? I would that that such an intercomparison study such as this would provide more concrete recommendations.

Response: The above results provide some directions for future model development, such as new heterogeneous or aqueous pathways for sulfate and nitrate formation under hazy condition, SOA chemical mechanism with new VOC precursors, yield data and approaches, and the dependence of aerosol optical properties on size distribution and mixing state. We have added this in the revised manuscript.

Line 788- 792: This sort of general conclusion about model inter-comparisons studies is rather tired. It could have been written based on results already in the literature and by speculation, without even conducting this inter-comparison study. I wish he authors could be more specific here regarding the findings specific to this study. The authors have not investigated all these possibilities and only barely scratched the surface at isolating and ascribing the uncertainties to specific processes.

Response: This paper presents model evaluation and intercomparison for meteorological fields, aerosol concentrations and optical properties with a series of observations over East Asia. While these models exhibit a generally good performance for PM_{2.5} concentration, the simulations of chemical compositions differ largely, which lead to the large differences in optical properties, such as AOD, and this would further affect direct and indirect aerosol effects, and consequently radiation and cloud in this region. We will compare and examine how the different AOD levels and aerosol properties affect radiation and cloud and explore the strength and affecting factors of aerosol-radiation-weather interactions in the companion paper part II.

Conclusion: It would be useful to have some text that looks forward to the next part of the paper. Will the authors be looking at evaluating the feedbacks, which were not examined in this study? Another reason I am looking for some more concrete explanations in this paper for why the models differ is that the situation will only become more complicated when examining feedback effects.

Responses: Thank you for the good suggestion. We briefly describe the objective of the next part of the paper in the conclusion: “This paper focused on the evaluation of the predictions of meteorological parameters and the predictions of aerosol mass, composition and optical depth. These factors play important roles in feedbacks impacting weather and climate through radiative and microphysical processes.”

Referee #2

The paper describes the setup of the MICS Asia Phase III model experiment. The results of seven simulations with online coupled meteorology-atmospheric chemistry models are shown and compared against observational data. The evaluation of online coupled air quality models and the outcome of the MICS-ASIA III model inter-comparison exercise are certainly worth to be published. However, the quality of the paper must be improved significantly before it can be published in ACP. In the first instance, the 'Results and discussions' section must be enhanced considerably. Attempts should be made to explain the reasons for the observed differences among the models. In particular, a more in depth discussion is necessary for those model results which look like outliers (for example solar radiation and ozone in Figure 5 and 6). Although the paper is overall well organized, it was nevertheless a tough read for me.

It is sometimes difficult to keep track of the different models and their respective setups.

Repetition of the model name along with the label once a while could improve this situation with only little effort. Enhancing the figure captions would also help. Finally, several sentences are quite convoluted and hard to understand. Splitting long sentences may improve the readability.

For these reasons and further reasons mentioned below, major revisions are required.

The specific comments include some suggestions how to improve the paper.

[Response: We appreciate the detailed and constructive comments from referee #2 to help improve our manuscript. We have added more explanations for the shown model differences in the revised version. We also add model names in figure captions, rewrite some sentences, and split long sentences to improve the readability. The detailed point-by-point responses and changes are listed below.](#)

Specific comments (also including minor points):

Abstract: In its current form the abstract raises the expectation that aerosol meteorology interactions will also be a topic of this paper. The reader may also expect that both episodes will be evaluated in the paper. Generally, the first part of the abstract is raising expectations, which are not fulfilled by the paper. Nevertheless, it is also OK to restrict the paper to the evaluation of first episode. However, the abstract must be reworded in this case in order to

avoid raising expectations which are not met later. It is not clear why the model evaluation is restricted to the first episode of the model inter-comparison and why the second episode is not considered.

Response: Thanks for pointing out. This manuscript is the overview paper for aerosol meteorology interactions topic, so we include it in the abstract to raise interests from audience in the companion papers. We deleted “Two winter months (January 2010 and January 2013) were selected as study periods, when severe haze occurred in North China.” in the abstract to avoid the expectation. We restricted to the first case because all models submitted results for January 2010 and less models for January 2013. For other topics within MICS-ASIA III, only year 2010 was simulated. Besides, more measurements data are available for year 2010. We also found that we had enough materials (mostly project overview) to present for current manuscript, so decided to move the evaluations of the second case into part II manuscript. We have removed all descriptions for 2013 in the manuscript.

Introduction: The authors should also consult publications describing related work, e.g. the model evaluation papers related to HTAP as well as AQMEII Phase 2 and Phase 3.

Response: It is a great idea to include information from HTAP and AQMEII projects. In the introduction, we add “Other ongoing related modeling frameworks include the Task Force on Hemispheric Transport of Air Pollution (TF HTAP) and the Air Quality Model Evaluation International Initiative (AQMEII). The TF HTAP was initiated to improve knowledge of the intercontinental or hemispheric transport and formation of air pollution, and its impacts on climate, ecosystems and human health (Galmarini et al., 2017; Huang et al., 2017). The AQMEII project specifically focuses on regional modeling domains over Europe and North America (Galmarini et al., 2017), within which aerosol meteorology interactions was studied (Forkel et al., 2015; Makar et al., 2015a, 2015b; San Jose et al., 2015) over Europe and North America.”

Line 106: ‘Various multi-scale models: Are there more than the five model and seven simulations described here?’

Response: Yes. This sentence briefly describes a bigger picture of MICS-ASIA Phase III, not

just Topic 3. Other topics include other models, such as GEOS-Chem. To avoid misunderstanding, I add one more sentence in the context: “A detailed overview of MICS-Asia Phase III, including descriptions of different research topics and participating models, will be published in a companion paper”.

Lines 114 –119: The references should not be restricted to Chinese authors.

Response: We include more references on this topic: “Forkel et al., 2015; Makar et al., 2015a, 2015b; San Jose et al., 2015” from AQMEII Phase 2.

Line 125: The paper claims to be ‘serving as the main repository of the information linked to Topic 3 simulations and comparisons.’ To achieve this aim, more details must be added to the descriptions of those models where the description is quite. Please add also some paragraphs in section 2.1 (similar to section 2.5) in order to make this section better readable.

Response: Sorry for the relatively poor presentation in section 2.1. We have reorganized section 2.1 following the format in section 2.5 to make it more readable.

Table 1: The numbers attached to ‘WRF-Chem’ and ‘NU-WRF’ are unnecessary and confusing. Therefore, they must be removed. On the other hand, the model version is an important information which must be added wherever applicable. The contents of the table are not precise: For example, for M1 ‘RACM’ must be replaced by ‘RACM-ESRL’ and ‘MADE’ by ‘MADE/VBS’. What climatological data are applied as boundary conditions for M6? Please consider to add also information about the details of the radiation calculation. As shown by Curci et al. (2015, Atmospheric Environment, Vol. 115) the inherent assumption have a strong influence on the calculated AODs.

Response: Thanks for these great suggestions. We delete the numbers attached to WRF-Chem and NU-WRF and add the model version. We replaced ‘RACM’ and ‘MADE’, and add climatological boundary data for each model. We add both longwave and shortwave parameterization information and mixing state information for each model to Table 1 and cite Curci et al. (2015) to emphasize the importance of radiation calculation on AOD.

Table S1: Please check also whether this table needs to be more specific (similar to Table 1). Please add also information about the two models which are missing in the table. Please consider also moving this table into the main part of the paper.

Response: To make it more specific, we add information for other missing models, including

climatological boundary data. We also merge S1 and Table 1.

Line 338: Please mention the source of the climatological data here or in Table 1.

Response: We add this in Table 1 in the revised manuscript.

Line 281: Please add the information which model uses the prescribed BVOC emissions and which one does the internal calculation (could also be added to Table 1).

Response: We add this in Table 1 in the revised manuscript.

Section 2.1: Please mention clearly in the text which aerosol meteorology interactions are switched on for each model.

Response: We include this information in Table 1 and add this in the revised manuscript.

Section 2.2: Please add some information on soil dust emissions (also to be included in Table 1).

Response: We add this information in Table 1 in the revised manuscript.

Section 2.3: Please include some information about the meteorological boundary conditions.

Response: We add this information in Table 1 in the revised manuscript.

Section 2.3: Please consider to add some information about the differences between the boundary values from the different data sources for selected variables (eventually to be included in the supplementary material). Looking into the boundary values may also help to understand differences between results of the different model runs.

Response: We add this information in the supplementary material as shown below. For aerosol species and gases except ozone, boundary conditions from the two models are pretty similar to each other. MOZART simulated ozone is higher than GEOS-Chem. Previous tests have showed that the influences of different boundary conditions have negligible impact on PM simulations, but larger impact on ozone (Abdallah et al., 2016). Topic 3 focuses on aerosol-weather-climate interactions in North China, so the impacts of different chemical boundary conditions are not quite important on our results. But it might be part of reason for poor ozone performance for some models. Thanks for this great suggestion. We have added this discussion in the revised paper.

Abdallah, C., Sartelet, K. and Afif, C., 2016. Influence of boundary conditions and anthropogenic emission inventories on simulated O₃ and PM_{2.5} concentrations over Lebanon. Atmospheric Pollution Research, 7(6), pp.971-979.

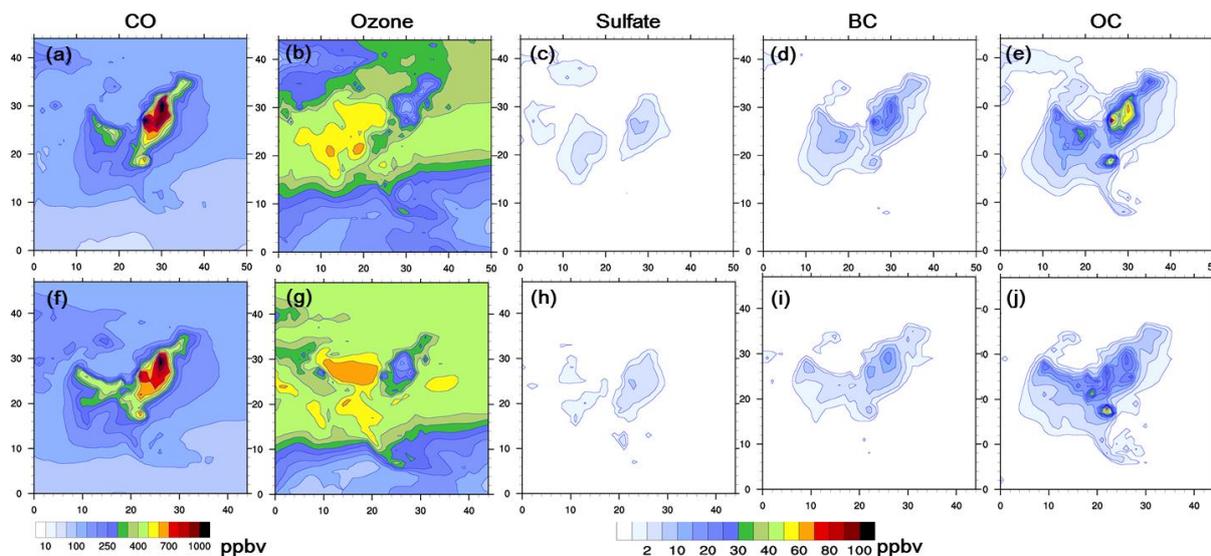


Figure Monthly mean near surface CO, ozone, sulfate, BC and OC from GEOS-Chem (a-e) and MOZART (f-i)

Line 389: This sentence is confusing. According to Table S1 all simulations considered here are performed with aerosol meteorology interactions switched on. Were the simulations additionally performed also without aerosol meteorology interactions for the investigation of feedback effects? Anyway, this could be mentioned in the introduction as well as in the outlook, but not at this place.

Response: Yes, Table S1 only specifies what kind of interactions were turned on. We also performed without interactions to check the differences. The differences between with and without interactions will be present in paper part II on aerosol-meteorology interactions. Following your suggestion, we move this sentence to the introduction section.

Line 392 –393: Why are abbreviations given for temperature, humidity, and wind, and units for the shortwave radiation?

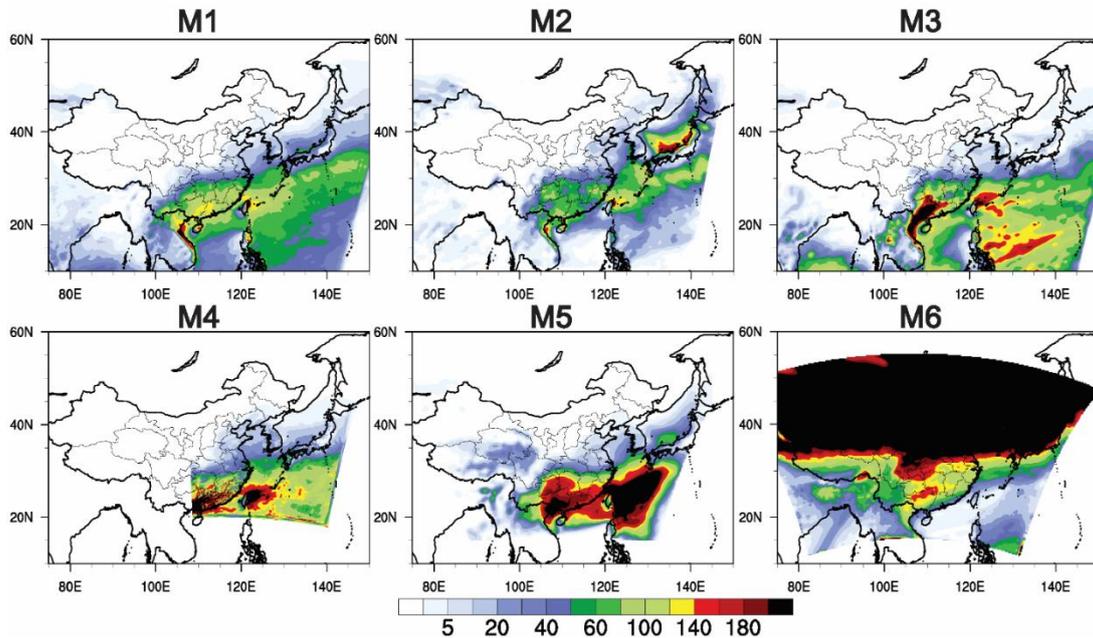
Response: To make it easier to show in figures.

Section 3: Why is the year 2013 described here, if it not discussed in the rest of the paper? Figure 3 could eventually be moved to the supplementary material.

Response: Following your suggestion, we move it to the supplementary material and remove corresponding sentences in the text.

Section 4.1: Cloud optical depth and integrated liquid water are important and should also be discussed (even, if no observational data are available).

Response: Following your suggestion, we add plots of integrated liquid water. We found this plot is good to explain why SWDOWN in north China is extremely low for M6. We added this point in the revised manuscript.



Line 427: This topic is not addressed in this paper.

Response: Thanks for pointing out. We delete it here and will provide discuss it the forthcoming companion paper.

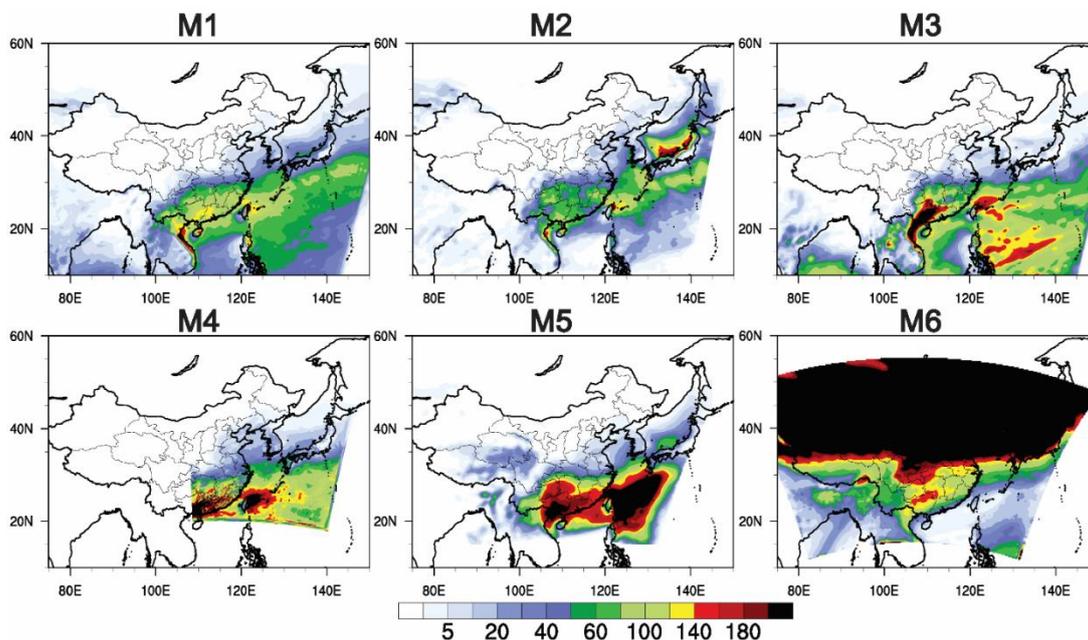
Line 473: What is the reason for the bad performance of M6 and M7? Please check also cloud cover. And why is this huge difference in radiation between M6 and M7 not reflected in T2? Why does M7, which is also WRF, show such a large difference to M1 and M2?

Response: The reason for M7 overestimation of SWDOWN in North China is the exclusion of aerosol-radiation interactions, which tends to overpredict SWDOWN especially in high PM days and areas. T2 looks fine because T2 nudging was used, so T2 is close to measurements in M7. M1 and M2 used very different settings, with aerosol-radiation interactions, and without using meteorology nudging. M6 simulates a lower SWDOWN due to overestimation of cloud integrated liquid water as mentioned above. Sorry for the confusion here. We add the descriptions of setting in Table 1 and text to make it clearer.

Line 480: Why is this the case? (Figure 5e), M6 and M7 show a better consistence with observations than over northern China sites.

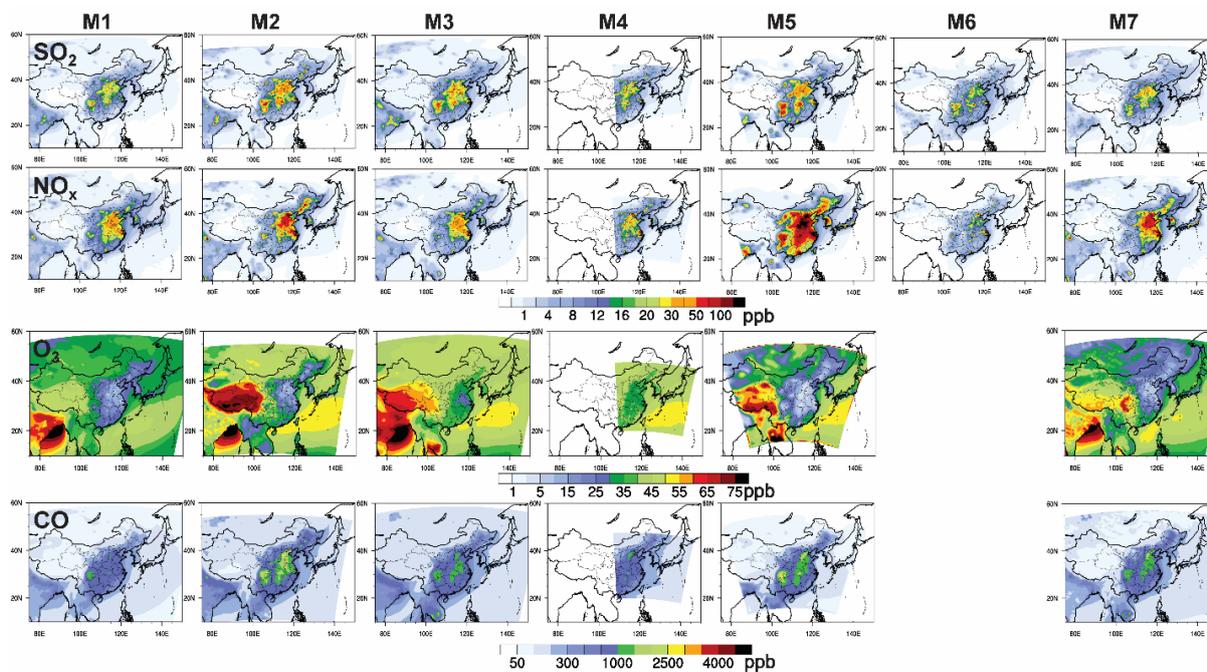
Response: We add plots of integrated liquid water, which may explain a better simulation of

M6 in southern China, and M7 shows a better agreement with observation because relatively lower PM level and weaker aerosol radiative effect in southern China.. We added this point in the revised manuscript.



Section 4.2: Spatial distributions of the gas phase pollutants (similar to figures 8 and 9) could be shown also in the supplement.

Response: Thanks for this suggestion. We have added in the revised manuscript.



Line 498: The ‘only’ is probably placed wrong here?

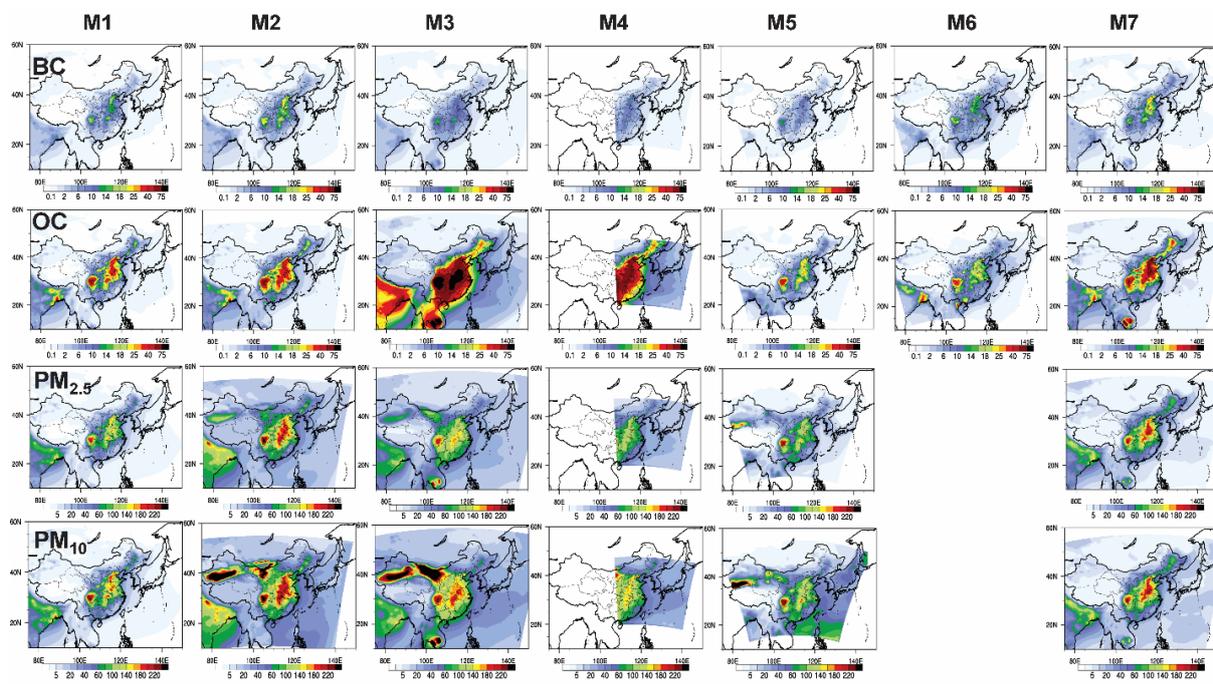
Response: Yes, we have removed it.

Line 504: Please give some evidence for this.

Response: This statement is based on the mean MBE averaged over all used CARE-China sites. It is typo here. We changed Beijing to “CARE-China sites”.

Line 521: Please give more evidence for this statement. And what is the contribution of soil dust to PM10? What is the contribution of different of the dust emission parameterizations?

Response: In winter of the north Huabei Plain, soil dust generally contributes about 10% to PM concentration, but there is also primary PM from anthropogenic activity, such as power plant, traffic, construction etc. and this part of PM mostly settles in coarse mode, which may contribute to PM₁₀, but it's not clear if all the models include this emission sector and this sector is of higher uncertainty compared with other anthropogenic emission sectors. The schemes of dust deflation are similar in WRF-Chem series and different in M5, M6 and M7, the discussion on different dust schemes is less important given the relatively small contribution to PM_{2.5}. However, we added the dust implementation in Table 1 to provide how it affect PM10, which are shown below. The implementation of wind-blown dust are mostly significant in northwestern regions, and less important in Beijing and surrounding regions.



Line 524 – 526: What is the reason for the negative correlation? On an hourly basis, the diurnal course of ozone should reflect the course of the solar radiation. Or are the correlations just calculated from on daily values? Please clarify. How do the diurnal courses look like for the

individual models? Could the parameterization of dry deposition or differences in the lateral boundary conditions or differences in biogenic VOC emissions explain the overestimation of simulated ozone for M3 and M4? It is surprising, that the overestimated ozone seems not to be related to solar radiation, so there are probably other reasons for this overestimation. How do the ozone profiles look like, can they contribute to an explanation?

Response: We think the major causes for O₃ overprediction in M3 and M4 are the combined effects of vertical diffusion and lateral boundary condition. M3 and M4 could predict larger vertical diffusivity coefficients, which leads to stronger vertical mixing, this not only results in lower NO_x concentration and weaker titration of O₃, but also stronger mixing of O₃ from lateral boundary down to near surface, consequently causing higher O₃ concentration than others.

Line 530: If the citation does not fit, it should be removed. Knote et al. (2015, Atmospheric Environment, Vol. 115) may probably suit better as this paper includes also a winter episode.

Response: Thanks. We replaced this citation with Knote et al. 2015.

Sections 4.3 and 4.4: It would be nice if soils dust were also included in the discussion.

Response: Thanks. We have added discussion of dust in the revised manuscript.

Lines 550 – 554: These general remarks are not necessary here.

Response: We remove this part.

Line 561: ‘M5 and M6 shows : : :’: Please observe the proper use of singular and plural (not only here but throughout the paper, e.g. line 582).

Response: We have changed these two and checked through the manuscript.

Lines 562 – 578: It is not clear in how far these statements apply to the models which are discussed here.

Response: We think these sentences are important to explain the possible reasons for the consistent underestimation of sulfate concentration from most of the models. Some important chemical processes for sulfate formation during hazy events are not implemented or fully considered in the participating models. M5 and M6 don't show the low biases possibly due to other compensatory processes, such as larger chemical reaction rate or lower deposition in the models.

Line 568 – 572: This sentence is quite hard to understand. Please split it into two or more sentences. Please also split other lengthy convoluted sentences.

Response: We have changed these two and checked through the manuscript.

Line 593: Is this just a general remark or is there some evidence for this? Does M7 really include heterogeneous nitrate formation?

Response: Yes, in default CMAQ, the heterogeneous nitrate related reactions include $\text{N}_2\text{O}_5 + \text{H}_2\text{O}$ and $2\text{NO}_2 + \text{H}_2\text{O}$

Line 608 and line 772: How can this be? Emissions were supposed to be the same for all models.

Response: We explain this in response to the same question from another reviewer. Emission should be similar among models, but treatments of deposition and aging process are different, we rewrite the sentence.

Line 609: Was the vertical distribution of the emissions not prescribed? If there are differences in the vertical distributions, they must be described.

Response: All the models release emission at the surface. We have rewritten this sentence.

Line 616: According to line 191, M2 does not include SOA formation

Response: We have removed M2 here.

Line 616 – 617: The statement about SORGAM does not fit here as M1 includes a VBS approach (Ahmadov, R., et al., 2012, J. Geophys. Res.).

Response: Sorry, we double check with the modeler, and they confirm that SORGAM instead of VBS is used. We have changed the description in Table 1 and in the text.

Line 619: ‘volatile’ seems to be missing here

Response: We changed it to semi-volatile VOCs

Lines 619 – 624: These lines include quite general statements. How are they related to the models discussed in this paper?

Response: We have deleted this sentence.

Lines 642 – 643: This should have been mentioned in section 2.2.

Response: We have added the dust parameterization in section 2.2.

Lines 669 – 672: Please avoid this kind of redundancies (not only here).

Response: We have removed it and checked through the paper.

Lines 677 – 680: What is the contribution of soil dust during these situations?

Line 688: Please split this sentence.

Response: We have splitted.

Line 690 – 691: This statement is true, but unnecessary.

Response: We have removed.

Lines 692 – 698: These are quite general statements, which must be related to the applied models.

Response: We have removed

Line 710: M2 does not have modes

Response: We added (size bins)

Line 712: What are the consequences of this? If necessary, this can be discussed based on the findings by Curci et al. (2015, Atmospheric Environment, Vol. 115).

Response: Thanks for this good suggestion. We added the conclusion from Curci et al. (2015) that external mixing can increase simulated AOD but core-shell assumption is a minor issue.

Figures and figure captions:

Figure captions in general: More detailed descriptions must be given in the figure captions (are daily values or hourly values shown, relevant area, etc., depending on what is shown in the figure).

Response: Thanks for pointing out. We have added these information.

Figure 1: It looks like the labels M3 and M4 in the legend of Figure 1 are mixed up (According to Table 1 and Figures 8 and 9 M4 is the small domain). Furthermore, using the same colors for the model domains as for the curves in Fig. 5 etc. would make the reading a bit easier.

Response: Thanks for pointing out. We noticed it after submission. We have changed it.

Caption of Figure 1: Repeating the model names in the caption (for example ‘M1: WRF-Chem, 45 km; M2 WRF-Chem, 50 km, M3 : : :’) would make the paper a bit ‘reader-friendly’.

Response: Thanks for this good suggestion. We have added: M1: WRF-Chem 45km; M2: WRF-Chem 50km; M3: NU WRF 45km; M4: NU-WRF 15km; M5: RIEMS-IAP 60km; RegCCMS 50km; WRF-CMAQ 45km

Caption of Figure 5: Please mention that this is a spatial average (for a, b and c: over which area) of daily values and explain error bars.

Response: Thanks for this good suggestion. We have added: (spatial daily values are averaged over measurements shown in S4 and S5; the error bars show the standard deviation of values over the measurement sites)

Figure 7: Please show also CO and PM2.5 although no measurements are available.

Response: Thanks for this good suggestion. We have changed.

Figures 8 and 9: Please include also PM10.

Response: Thanks for this good suggestion. We have changed.

Figures 8 and 9: The split into two figures appears quite arbitrary. Perhaps it would look better if the figures are organized differently: One figure with 7 rows (M1 – M7) and 3 columns (PM10, PM2.5, BC) and one figure with 7 rows (M1 – M7) and 4 columns (SO24, NO3, NH+4, OC) – no obligation to do this, just a suggestion.

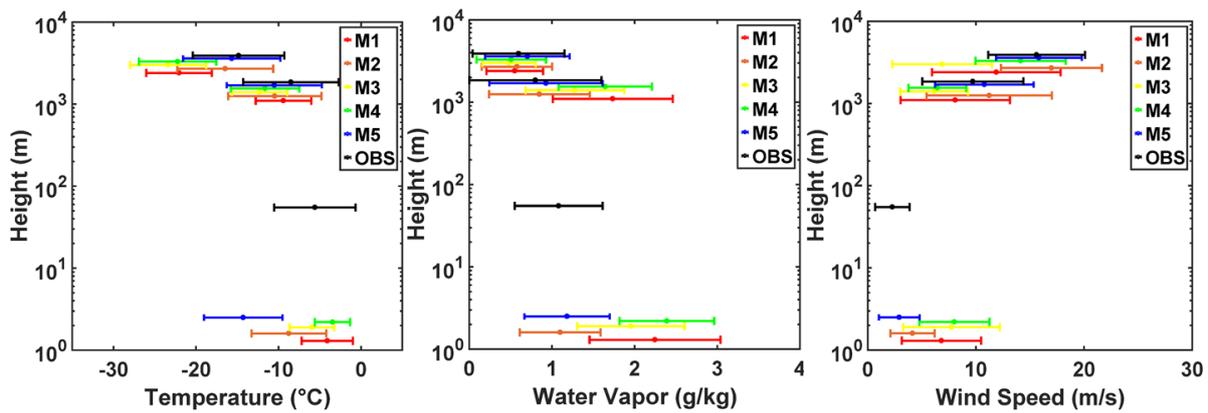
Response: Thanks for this good suggestion. We have added PM10 in Figure 9.

Spatial distributions of the gas phase pollutants (similar to Figures 8 and 9) would be nice in the supplement.

Response: Thanks for this good suggestion. We have added gas pollutants.

Figure S4: This figure seems to be contorted. Please improve the quality. Why are all the lines within the single ‘height groups’ (e.g. at ‘1 km’) at different heights? Please explain in the figure caption. Also: explain what is shown here (daily values, hourly values, : : :?)

Response: Do you mean Figure S7? I used slight differences to better show there values, otherwise they overlap with each other. The shown values are monthly mean. I have added these information (near surface observation is at 55m and model predictions are at 2m; comparisons are conducted at near surface, 1km and 3km; shifts in heights are made to make it clearer to avoid overlapping) in the revised caption and replotted it to sole the contortion problem.



1 **Air Quality and Climate Change, Topic 3 of the Model Inter-Comparison**
2 **Study for Asia Phase III (MICS-Asia III), Part I: overview and model**
3 **evaluation**

4 Meng Gao^{1,2}, Zhiwei Han^{3,4}, Zirui Liu⁵, Meng Li^{6,13}, Jinyuan Xin⁵, Zhining Tao^{7,8}, Jiawei Li⁴, Jeong-Eon
5 Kang⁹, Kan Huang¹⁰, Xinyi Dong¹⁰, Bingliang Zhuang¹¹, Shu Li¹¹, Baozhu Ge⁵, Qizhong Wu¹², Yafang
6 Cheng¹³, Yuesi Wang⁵, Hyo-Jung Lee⁹, Cheol-Hee Kim⁹, Joshua S. Fu¹⁰, Tijian Wang¹¹, Mian Chin⁸,
7 Jung-Hun Woo¹⁴, Qiang Zhang⁶, Zifa Wang^{4,5}, Gregory R. Carmichael¹

8 1 Center for Global and Regional Environmental Research, University of Iowa, Iowa City, IA, USA

9 2 John A. Paulson School of Engineering and Applied Sciences, Harvard University, Cambridge, MA, USA

10 3 Key Laboratory of Regional Climate-Environment for Temperate East Asia, Institute of Atmospheric Physics,
11 Chinese Academy of Sciences, Beijing, China

12 4 University of Chinese Academy of Sciences, Beijing 100049, China

13 5 State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric Chemistry, Institute of
14 Atmospheric Physics, Chinese Academy of Sciences, Beijing, China

15 6 Ministry of Education Key Laboratory for Earth System Modeling, Center for Earth System Science, Tsinghua
16 University, Beijing, China

17 7 Universities Space Research Association, Columbia, MD, USA

18 8 NASA Goddard Space Flight Center, Greenbelt, MD, USA

19 9 Department of Atmospheric Sciences, Pusan National University, Busan, South Korea

20 10 Department of Civil and Environmental Engineering, University of Tennessee, Knoxville, TN, USA

21 11 School of Atmospheric Sciences, Nanjing University, Nanjing, China

22 12 College of Global Change and Earth System Science, Beijing Normal University, Beijing, China

23 13 Multiphase Chemistry Department, Max Planck Institute for Chemistry, Mainz, Germany

24 14 Department of Advanced Technology Fusion, Konkuk University, Seoul, South Korea

25 Correspondence to: M. Gao (mgao2@seas.harvard.edu), Z. Han (hzw@mail.iap.ac.cn), and G. R.

26 Carmichael (gcarmich@engineering.uiowa.edu)

27

28 **Abstract**

29 Topic 3 of the Model Inter-Comparison Study for Asia (MICS-Asia) Phase III examines how
30 online coupled air quality models perform in simulating high aerosol pollution in the North
31 China Plain region during wintertime haze events and evaluates the importance of aerosol
32 radiative and microphysical feedbacks. A comprehensive overview of the MICS-ASIA III Topic
33 3 study design, including descriptions of participating models and model inputs, the experimental
34 designs, and results of model evaluation, are presented. ~~Two winter months (January 2010 and~~
35 ~~January 2013) were selected as study periods, when severe haze occurred in North China.~~
36 ~~Simulations were designed to evaluate radiative and microphysical feedbacks, together and~~
37 ~~separately, relative to simulations without feedbacks.~~ Six modeling groups from China, Korea
38 and the United States submitted results from seven applications of online coupled chemistry-
39 meteorology models. Results are compared to meteorology and air quality measurements,
40 including the Campaign on Atmospheric Aerosol Research Network of China (CARE-China)
41 network, and the Acid Deposition Monitoring Network in East Asia (EANET). ~~The analysis~~
42 ~~focuses on model evaluations and aerosol effects on meteorology and air quality, and potentially~~
43 ~~other interesting topics, such as the impacts of model resolutions on aerosol radiation weather~~

44 ~~interactions. The model evaluations for January 2010 show that current online-coupled~~
45 ~~meteorology-chemistry model can generally well reproduced meteorological features and~~
46 ~~variations of major air pollutants, including aerosol concentrations. The correlation coefficients~~
47 ~~between multi-model ensemble mean and observed near-surface temperature, water vapor~~
48 ~~mixing ratio and wind speeds can reach as high as 0.99, 0.99 and 0.98.~~The correlation
49 coefficients between multi-model ensemble mean and the CARE-China observed near-surface
50 air pollutants range from 0.51 to 0.94 (0.51 for ozone and 0.94 for PM_{2.5}) for January 2010.
51 However, large discrepancies exist between simulated aerosol chemical compositions from
52 different models, which is due to different parameterizations of chemical reactions. The
53 coefficient of variation (standard deviation divided by ~~average~~the mean) can reach above 1.3 for
54 sulfate in Beijing, and above 1.6 for nitrate and organic aerosol in coastal regions, indicating
55 these compositions are less consistent from different models. During clean periods, simulated
56 Aerosol Optical Depths (AOD) from different models are ~~consistent~~similar, but peak values
57 differ during severe haze events, which can be explained by the differences in simulated
58 inorganic aerosol concentrations and the hygroscopic growth efficiency (affected by varied RH).
59 These results ~~provide some brief senses of~~present how current online-coupled meteorology-
60 chemistry models reproduce severe haze events, and provide some directions for future model
61 improvements, such as new heterogeneous or aqueous pathways for sulfate and nitrate
62 formation under hazy condition, secondary organic aerosol (SOA) formation chemical
63 mechanism with new volatile organic compounds (VOCs) precursors, yield data and approaches,
64 and the dependence of aerosol optical properties on size distribution and mixing state.

65

66 **1 Introduction**

67 Air pollution in Asia, particularly in China and India, has been an increasingly important research
68 topic, and has attracted enormous media coverage since about 60% of the world population live
69 and are exposed to extremely unhealthy air in this region. It is estimated that outdoor air
70 pollution brings about 3.3 million premature deaths per year worldwide ~~but with most deaths~~
71 ~~occur~~ primarily in Asia (Lelieveld et al., 2015). In addition, the impacts of regional and
72 intercontinental transport of Asian pollutants on air quality and climate change have been
73 frequently reported (Akimoto, 2003; Menon et al., 2002, Ramanathan and Carmichael, 2008).
74 Chemical transport models have been developed and applied to study various air pollution issues
75 in Asia. For example, an Eulerian regional scale acid deposition and photochemical oxidant
76 model was developed in the United States (Carmichael and Peters, 1984; Carmichael et al., 1986;
77 Carmichael et al., 1991) and applied to study long-range transport of sulfur oxides (SO_x), dust
78 and ozone production in East Asia (Carmichael et al., 1998; Xiao et al., 1997). ~~;-). a~~ A nested
79 urban and regional scale air quality prediction modeling system was developed and applied to
80 investigate ozone pollution in Taiwan (Wang et al., 2001). Although important advances have
81 taken place in air quality modeling, large uncertainties still remain, which are related to
82 inaccurate and/or incomplete emission inventories, poorly represented initial and boundary
83 conditions and missing or poorly parameterized physical and chemical processes (Carmichael et
84 al., 2008a).

85 Furthermore, many models used to study air quality in Asia have been developed in other regions
86 (e.g., USA and Europe), and the assumptions and parameterizations included in these models
87 may not be applicable to the Asian environment. In order to develop a common understanding of

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

88 model performance and uncertainties in Asia, and to further develop the models for Asian
89 applications, a model inter-comparison study was initiated, i.e., Model Inter-Comparison Study
90 for Asia Phase I (MICS-Asia I), in 1998 during a workshop on Transport of Air Pollutants in
91 Asia in Austria. The focus of MICS-Asia Phase I was to study long-range transport and
92 deposition of sulfur within Asia in support of on-going acid deposition studies. Eight long-range
93 transport models from six institutes in Korea, Japan, Denmark, the USA, and Sweden
94 participated in MICS-Asia I. Multi-model results of sulfur dioxide (SO₂) and sulfate
95 concentrations, and wet deposition amounts in January and May 1993 were compared with
96 surface observations in East Asia (Carmichael et al., 2002). Source-receptor relationships and
97 how model structure and parameters affect model performance were also discussed during this
98 phase (Carmichael et al., 2002). In 2003, MICS-Asia Phase II was initiated to include more
99 species, including nitrogen compounds, ozone and aerosols. The study period was expanded to
100 cover two different years and three different seasons, and global inflow to the study domain was
101 also considered (Carmichael et al., 2008b). Nine modeling groups from Korea, Hong Kong,
102 Japan, the USA, Sweden, and France participated in this phase. Seven topics (i.e., ozone and
103 related precursors, aerosols, acid deposition, global inflow of pollutants and precursors to Asia,
104 model sensitivities to aerosol parameterization, analysis of emission fields, and detailed analyses
105 of individual models) were discussed and published in a special issue of Atmospheric
106 Environment (Carmichael et al., 2008b).
107 In 2010, MICS-Asia phase III was launched and three topics for this phase were decided during
108 the first and second Workshop on Atmospheric Modeling in East Asia. Phase III aims to evaluate
109 strengths and weaknesses of current air quality models and provide techniques to reduce
110 uncertainty in Asia (Topic 1), to develop a reliable anthropogenic emission inventory in Asia

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

111 (Topic 2), and to evaluate aerosol-weather-climate interactions (Topic 3). Various multi-scale
112 models participated in this phase and the study periods range from year to month depending on
113 study topics. This phase uses data from the Acid Deposition Monitoring Network in East Asia
114 (EANET), in addition to new observations related to atmospheric chemistry in the region. [A](#)
115 [detailed overview of MICS-Asia Phase III, including descriptions of different research topics and](#)
116 [participating models, will be published in a companion paper.](#) An important advance to this
117 phase is the inclusion of multiple online-coupled chemistry-meteorology models to investigate
118 aerosol-weather-climate interactions, which is the target of topic 3. On-line coupled models are
119 playing important roles in air quality, meteorology and climate applications, but many important
120 research questions remain (Baklanov et al., 2017).

Formatted: Highlight

121 The influences of aerosols on meteorology, e.g., radiation, temperature, boundary layer heights,
122 winds, etc. and PM_{2.5} concentrations have been examined in previous studies using different
123 online coupled models (Forkel et al., 2015; Gao et al., 2016a, 2016b, 2017a, 2017b; Han et al.,
124 2012, 2013; Makar et al., 2015a, 2015b; San Jose et al., 2015; Tao et al., 2015, 2016; Wang et
125 al., 2014; Zhang et al., 2010). In general, there are two ways of online coupling: online integrated
126 coupling (meteorology and chemistry are simulated using the same model grid, and one main
127 time step is used to integrate) and online access coupling (meteorology and chemistry are
128 independent but data are exchanged on a regular basis) (Baklanov et al., 2014). These two
129 different coupling ways can lead to uncertainties in the results of aerosol-weather-climate
130 interactions. Even using the same coupling way, different parameterizations in different online
131 models causes uncertainties as well. Thus, it is important to inter-compare how different online
132 models simulate aerosol-weather-climate interactions-, [particularly in heavily polluted Asian](#)
133 [region. Other ongoing related modeling frameworks include the Task Force on Hemispheric](#)

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

134 Transport of Air Pollution (TF HTAP) and the Air Quality Model Evaluation International
135 Initiative (AQMEII). The TF HTAP was initiated to improve knowledge of the intercontinental
136 or hemispheric transport and formation of air pollution, and its impacts on climate, ecosystems
137 and human health (Galmarini et al., 2017; Huang et al., 2017). The AQMEII project specifically
138 focuses on regional modeling domains over Europe and North America (Galmarini et al., 2017),
139 within which aerosol meteorology interactions was studied (Forkel et al., 2015; Makar et al.,
140 2015a, 2015b; San Jose et al., 2015) over Europe and North America.

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

141 This paper presents and overview of the MICS-ASIA III Topic 3, serving as the main repository
142 of the information linked to Topic 3 simulations and comparisons. Specifically, this paper aims
143 to archive the information of participating models, how the experiments, and results of model
144 evaluation. ~~Simulations were designed to evaluate radiative and microphysical feedbacks,~~
145 ~~together and separately, relative to simulations without feedbacks.~~ The results of the MICS-Asia
146 Topic 3 experiments looking at the direct and indirect effects during heavy haze events will be
147 published in a companion paper, part II. This paper is organized as follows: in Section 2, we
148 provide the inter-comparison framework of Topic 3, including the participating models,
149 emissions, boundary conditions, observational data, and analysis methodology. Section 3
150 presents the general descriptions of the study periods and Section 4 presents comparisons and
151 discussions focused on the results related to the meteorological and air pollution conditions
152 during the January 2010 heavy haze episode. The results of January 2013 haze episode and
153 detailed analysis of the direct and indirect effects will be presented in a companion paper.

154 **2 Inter-comparison framework**

155 In North China, severe aerosol pollution frequently happens and attracts enormous interests from
156 both public and scientific communities (Cheng et al., 2016; Gao et al., 2015, 2016a, 2016b,
157 2016c, 2017). Two winter months in which severe haze episodes happened in North China were
158 selected as the study periods for Topic 3. During these two months, maximum hourly PM_{2.5}
159 concentration in urban Beijing reached ~500 µg/m³ and 1000µg/m³, respectively. Compared to
160 the China Grade 1 24-h PM_{2.5} standard (35µg/m³), daily mean PM_{2.5} concentrations in urban
161 Beijing exceeded this standard for 20 days and 27 days within these two months, respectively.
162 The dramatically high aerosol loadings during these two hazy months substantially affected
163 radiation transfer, and provide a good opportunity to study the aerosol effects on weather, air
164 quality and climate. In this study, the participants were required to use common emissions to
165 predict air quality during these two months and submit requested model variables. The emissions
166 were placed on a publicly accessible website. Six modeling groups submitted results for Topic 3.
167 In this section, we briefly describe these models and their configurations, introduce the emission
168 inventories (including anthropogenic, biogenic, biomass burning, air and ship, and volcano
169 emissions), observational datasets, and describe the analysis methodology.

Formatted: Highlight

170 2.1 Participating models

171 Table 1 summarizes the characteristics of the participating models. These models include: one
172 application of the Weather Research Forecasting model coupled with Chemistry (WRF-Chem,
173 Fast et al., 2006; Grell et al., 2005) by Pusan National University (PNU) (M1), one application of
174 the WRF-Chem model by the University of Iowa (UIOWA) (M2), two applications (two
175 domains: 45km and 15km horizontal resolutions) of the National Aeronautics and Space
176 Administration (NASA) Unified WRF (NU-WRF, Peters-Lidard et al., 2015; Tao et al., 2013)
177 model by the Universities Space Research Association (USRA) and NASA's Goddard Space

Formatted: Highlight

Formatted: Highlight

178 Flight Center (M3 and M4), one application of the Regional Integrated Environment Modeling
179 System with Chemistry (RIEMS-Chem, Han et al., 2010) by the Institute of Atmospheric Physics
180 (IAP), Chinese Academy of Sciences (M5), one application of the coupled Regional Climate
181 Chemistry Modeling System (RegCCMS, Wang et al., 2010) from Nanjing University (M6), and
182 one application of the coupled WRF-CMAQ (Community Multiscale Air Quality) model by the
183 University of Tennessee at Knoxville (UTK) (M7). These models are all online coupled, which
184 enables aerosol-weather-climate interactions. Domain setting of each model application is shown
185 in Figure 1. The domains of M2, M5, and M6 (UIOWA, IAP, and NJU in Figure 1) cover most
186 areas of East Asia, including China, North Korea, South Korea, Japan, Mongolia, and north parts
187 of Southeast Asia. M1, M3 and M7 domains (PNU, NASA D01 and UTK) include more
188 countries in Southeast and South Asia. M4 (NASA D02) covers east China, Korea and Japan.

Formatted: Highlight

Formatted: Highlight

189 The descriptions of major model settings are listed below. More descriptions including
190 microphysics, radiation, and boundary layer, are listed in Table 1.

191 (1) Model grids: The horizontal model resolutions of these applications range from 15km to
192 60km (Table 1). Model vertical resolutions vary from 16 to 60 layers (Table 1) and the set model
193 top pressures range from 100mb to 20mb (Table 1).

194 (2) Gas phase chemistry: At PNU (M1), the RACM-ESRL (Regional Atmospheric Chemistry
195 Mechanism, Earth System Research Laboratory) gas phase chemistry was used. RACM was
196 developed based on Regional Acid Deposition Model (RADM2) to simulate regional
197 atmospheric chemistry (Stockwell et al., 1997) (including 237 reactions) and the rate coefficients

Formatted: Highlight

198 were updated in RACM ESRL version (Kim et al., 2009). At the University of Iowa (M2),

Formatted: Highlight

199 CBMZ (Carbon-Bond Mechanism version Z) gas phase chemistry was used. CBMZ (Zaveri and
200 Peters, 1999) extends the original CBM4 mechanism to function properly at larger spatial and

Formatted: Highlight

201 longer timescales. The augmented CBM4 scheme includes 67 species and 164 reactions. The
202 NU-WRF model (M3 and M4) uses RADM2 for gas phase chemistry. Both the RIEMS-Chem
203 model from IAP (M5) and the RegCCMS model from NJU (M6) used CBM4 to calculate gas
204 phase chemistry (Gery et al., 1989). The CBM4 version incorporated in RIEMS-Chem (M5)
205 includes 37 species and 91 reactions. The version of CBM4 implemented in RegCCMS (M6)
206 consists of 36 reactions (4 photolysis reactions) and 20 species (Wang et al., 2010). M7 applied
207 SAPRC 99 to simulate gas phase chemistry. The SAPRC99 mechanism implanted within the
208 CMAQ model has 88 species and 213 chemical reactions (Carter, 2000a, b).

209 (3) Aerosol modules: MADE/SORGAM (Modal Aerosol Dynamics Model for
210 Europe/Secondary Organic Aerosol Model) aerosol module was coupled and used in M1. MADE
211 uses 3 log-normal modes (Aitken, accumulation, coarse) and simulates major aerosol
212 compositions, including sulfate, ammonium, nitrate, sea-salt, black carbon (BC), and organic
213 carbon (OC). M2 uses 8 bin MOSAIC (Model for Simulating Aerosol Interactions and
214 Chemistry) aerosol module. MOSAIC considers major aerosol species at urban, regional and
215 global scales, including sulfate, nitrate, ammonium, sodium, chloride, EC, and other unspecified
216 inorganic species (such as inert minerals, trace metals, and silica) (Zaveri et al. 2008). The
217 MOSAIC version used in M2 includes some aqueous reactions but no SOA formation. At
218 NASA, the GOCART aerosol model (Chin et al., 2002) was coupled to RADM2 gas phase
219 chemistry, and incorporated into the NU-WRF model (M3 and M4) to simulate major
220 tropospheric aerosol species, including sulfate, BC, OC, dust, and sea-salt. In this aerosol model,
221 10% of organic compounds from the volatile organic compounds (VOCs) emission inventory are
222 assumed to be converted to SOA (Chin et al., 2002). Aerosols in RIEMS-Chem include sulfate,
223 nitrate, ammonium, BC, OC, SOA, 5 bins of soil dust, and 5 bins of sea salt (Han et al., 2012).

Formatted: Highlight

224 ISORROPIA (Nenes et al., 1998) is coupled to RIEMS-Chem to treat thermodynamic
225 equilibrium process and to simulate inorganic aerosols. SOA production from primary
226 anthropogenic and biogenic VOCs is calculated using a bulk aerosol yield method according to
227 Lack et al. (2004). RegCCMS also used ISORROPIA to calculate inorganic aerosols (Wang et
228 al., 2010). For implementation of aerosol effects, sulfate radiative properties were treated
229 following Kiehl and Briegleb (1993). OC were assumed to have the same properties as sulfate,
230 and the wavelength-dependent radiative properties of BC follows Jacobson (2001). AE6 aerosol
231 (the sixth-generation CMAQ aerosol module) mechanism is coupled with WRF. Compared to
232 previous version of CMAQ aerosol modules, AE6 improves SOA treatments, adds a new
233 heterogeneous N₂O₅ hydrolysis parameterization and ads a new gas-to-particle mass transfer for
234 coarse aerosols in sea-salt emissions (Yu et al., 2014). There are seven components including
235 water soluble mass, water insoluble mass, elemental carbon, sea salt, water, diameters and
236 standard deviations passed to WRF to directly change radiation calculations.

237 (4) Meteorological boundary and initial conditions: M1, M2, M5 and M7 use the National
238 Centers for Environmental Prediction (NCEP) final analysis (FNL) data to drive the model; M3
239 and M4 use NASA MERRA reanalysis data and M6 uses NCEP-NCAR reanalysis 1 dataset.

240 (5) Soil dust: M1, M6 and M7 do not include soil dust calculation. M3 and M4 use GOCART
241 dust module, and M2 uses a GOCART version that modified by AFWA (Air Force Weather
242 Agency). M5 uses a dust module that described in Han et al. (2004).

243 (6) Mixing state: M6 assumes external mixing, while other models use internal mixing
244 treatments.

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

Formatted: Subscript

Formatted: Subscript

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

245 Many previous studies have underscored that the choice of gas phase mechanism and aerosol
246 models are of great importance for simulating air pollutants (Knote et al., 2015). The different
247 gas phase chemistry and aerosol modules used in the participating models are expected to yield
248 notable differences in performances, which are shown later in section 4.

Formatted: Highlight

249 The horizontal model resolutions of these applications range from 15km to 60km (Table 1).
250 Model vertical resolutions vary from 16 to 60 layers (Table 1) and the set model top pressures
251 range from 100mb to 20mb.

252 Gas phase chemistry and aerosol modules are key components of chemical transport models.
253 Although the WRF-Chem and NU-WRF models were applied at three institutes (PNU, UIOWA,
254 and NASA), different gas phase chemistry and aerosol modules were used. At PNU (M1), the
255 RACM-ESRL (Regional Atmospheric Chemistry Mechanism, Earth System Research
256 Laboratory) gas phase chemistry coupled with MADE/VBS (Modal Aerosol Dynamics Model
257 for Europe/Volatility Basis set) aerosol module was used. RACM was developed based on
258 Regional Acid Deposition Model (RADM2) to simulate regional atmospheric chemistry
259 (Stockwell et al., 1997) (including 237 reactions) and the rate coefficients were updated in
260 RACM-ESRL version (Kim et al., 2009). MADE uses 3 log-normal modes (Aitken,
261 accumulation, coarse) and simulates major aerosol compositions, including sulfate, ammonium,
262 nitrate, sea-salt, black carbon (BC), and organic carbon (OC). In addition, the VBS method was
263 implemented to simulate secondary organic aerosols (SOA). At the University of Iowa (M2),
264 CBMZ (Carbon Bond Mechanism version Z) gas phase chemistry coupled with an 8 bin
265 MOSAIC (Model for Simulating Aerosol Interactions and Chemistry) aerosol module was
266 applied. CBMZ (Zaveri and Peters, 1999) extends the original CBM4 mechanism to function
267 properly at larger spatial and longer timescales. The augmented CBMZ scheme includes 67

268 species and 164 reactions. MOSAIC considers major aerosol species at urban, regional and
269 global scales, including sulfate, nitrate, ammonium, sodium, chloride, EC, and other unspecified
270 inorganic species (such as inert minerals, trace metals, and silica) (Zaveri et al. 2008). MOSAIC
271 includes some aqueous reactions but no SOA formation. At NASA, the GOCART aerosol model
272 (Chin et al., 2002) was coupled to RADM2 gas phase chemistry, and incorporated into the NU-
273 WRF model (M3 and M4) to simulate major tropospheric aerosol species, including sulfate, BC,
274 OC, dust, and sea salt. In this aerosol model, 10% of organic compounds from the volatile
275 organic compounds (VOCs) emission inventory are assumed to be converted to SOA (Chin et al.,
276 2002). Both the RIEMS-Chem model from IAP (M5) and the RegCCMS model from NJU (M6)
277 used CBM4 to calculate gas phase chemistry (Gery et al., 1989). The CBM4 version
278 incorporated in RIEMS-Chem (M5) includes 37 species and 91 reactions. The version of CBM4
279 implemented in RegCCMS (M6) consists of 36 reactions (4 photolysis reactions) and 20 species
280 (Wang et al., 2010). M7 applied SAPRC-99 coupled to the sixth generation
281 CMAQ aerosol module (AE6) to simulate gas phase chemistry and aerosol formation. The
282 SAPRC99 mechanism implanted within the CMAQ model has 88 species and 213 chemical
283 reactions (Carter, 2000a,b). At NASA, the GOCART aerosol model (Chin et al., 2002) was
284 coupled to RADM2 gas phase chemistry, and incorporated into the NU-WRF model (M3 and
285 M4) to simulate major tropospheric aerosol species, including sulfate, BC, OC, dust, and sea-
286 salt. In this aerosol model, 10% of organic compounds from the volatile organic compounds
287 (VOCs) emission inventory are assumed to be converted to SOA (Chin et al., 2002).
288
289 Both the RIEMS-Chem model from IAP (M5) and the RegCCMS model from NJU (M6) used
290 CBM4 to calculate gas phase chemistry (Gery et al., 1989). The CBM4 version incorporated in

291 ~~RIEMS-Chem (M5) includes 37 species and 91 reactions, and aerosols in RIEMS-Chem include~~
292 ~~sulfate, nitrate, ammonium, BC, OC, SOA, 5 bins of soil dust, and 5 bins of sea salt (Han et al.,~~
293 ~~2012). ISORROPIA (Nenes et al., 1998) is coupled to RIEMS-Chem to treat thermodynamic~~
294 ~~equilibrium process and to simulate inorganic aerosols. SOA production from primary~~
295 ~~anthropogenic and biogenic VOCs is calculated using a bulk aerosol yield method according to~~
296 ~~Lack et al. (2004). A lognormal size distribution is assumed for inorganic aerosols, BC, and OC,~~
297 ~~with median radius of 0.07 μm , 0.01 μm , and 0.02 μm , and geometric standard deviation of 2.0,~~
298 ~~2.0, and 2.2, respectively. The schemes for soil dust deflation and sea salt generation were from~~
299 ~~Han et al. (2004), which used 5 size bins (0.1–1.0, 1.0–2.0, 2.0–4.0, 4.0–8.0, 8.0–20.0 μm) to~~
300 ~~represent dust and sea salt size distribution. The refractive indices of aerosol components were~~
301 ~~mainly derived from the OPAC (Optical Properties of Aerosols and Clouds) database. Aerosol~~
302 ~~extinction coefficient as well as single scattering albedo and asymmetry factor are calculated by a~~
303 ~~Mie theory-based parameterization developed by Ghan and Zaveri (2007), which has a high~~
304 ~~computational efficiency with similar degree of accuracy compared with complete Mie code. An~~
305 ~~internal mixture of aerosols was assumed in this region of large emissions. A method known as~~
306 ~~kappa (κ) parameterization (Petters and Kreidenweis, 2007) was adopted to represent the aerosol~~
307 ~~hygroscopic growth.~~

308 ~~The version of CBM4 implemented in RegCCMS (M6) consists of 36 reactions (4 photolysis~~
309 ~~reactions) and 20 species (Wang et al., 2010). RegCCMS also used ISORROPIA to calculate~~
310 ~~inorganic aerosols (Wang et al., 2010). For implementation of aerosol effects, sulfate radiative~~
311 ~~properties were treated following Kiehl and Briegleb (1993), OC were assumed to have the same~~
312 ~~properties as sulfate, and the wavelength-dependent radiative properties of BC follows Jacobson~~
313 ~~(2001).~~

314 ~~M7 applied SAPRC 09 coupled to the sixth generation CMAQ aerosol module (AE6) to simulate~~
315 ~~gas phase chemistry and aerosol formation. The SAPRC09 mechanism implanted within the~~
316 ~~CMAQ model has 88 species and 213 chemical reactions (Carter, 2000a,b). AE6 aerosol~~
317 ~~mechanism is to couple with WRF. There are seven components including water soluble mass,~~
318 ~~water insoluble mass, elemental carbon, sea salt, water, diameters and standard deviations passed~~
319 ~~to WRF. Many previous studies have underscored that the choice of gas phase mechanism and~~
320 ~~aerosol models are of great importance for simulating air pollutants (Knote et al., 2015; Zhang et~~
321 ~~al., 2012). The different gas phase chemistry and aerosol modules used in the participating~~
322 ~~models are expected to yield notable differences in performances, which are shown later in~~
323 ~~section 4.~~

324 **2.2 Emissions**

325 The accuracy of air quality modeling results highly depends on the quality and reliability of
326 emission inventory. Accordingly, a new Asian emission inventory was developed for MICS-III
327 by integrating state-of-the-art national/regional inventories to support this model inter-
328 comparison study (Li et al., 2017). This is the major theme of MICS-ASIA III Topic 2. These
329 emissions, along with biogenic emissions, biomass burning emissions, emissions from air and
330 ship, and volcano emissions were used. This section offers some basic descriptions of these
331 provided emissions.

332 **2.2.1 Anthropogenic emissions**

333 The state-of-the-art anthropogenic emission inventory for Asia (MIX) was developed by
334 incorporating five inventories, including the REAS inventory for Asia developed at the Japan
335 National Institute for Environmental Studies (NIES), the MEIC inventory for China developed at

Formatted: Highlight

336 Tsinghua University, the high resolution ammonia (NH₃) emission inventory in China developed
337 at Peking University, the Indian emission inventory developed at Argonne National Laboratory
338 in the United States, and the CAPSS Korean emission inventory developed at Konkuk University
339 (Li et al., 2017). This MIX inventory includes emissions for ten species, namely SO₂, nitrogen
340 oxides (NO_x), carbon monoxide (CO), non-methane volatile organic compounds (NMVOC),
341 NH₃, PM₁₀, PM_{2.5}, BC, OC, and carbon dioxide (CO₂). NMVOC are provided with CB-05 and
342 SAPRC-99 speciation datasets. Emissions of these species were prepared for years 2008 and
343 2010 in monthly temporal resolution and 0.25 degree spatial resolution. Weekly/diurnal profiles
344 were also provided. Five sectors were considered, namely industry, power generation, residential
345 sources, transportation and agriculture. Figure 2 shows the spatial maps of these ten species for
346 January 2010. Emissions of most of these species exhibit similar spatial patterns, with enhanced
347 values in east China and lower values in north and south India. Emissions of NH₃ display a
348 different spatial distribution, with pronounced values in India and lower values in north China
349 (Figure 2). More detailed description of this emission inventory is documented in Li et al.
350 (2017).

Formatted: Highlight

351 2.2.2 Biogenic emissions

352 Terrestrial ecosystems generate miscellaneous-various chemical species, including volatile and
353 semi-volatile compounds, which play important roles in atmospheric chemistry and are the
354 largest contributor to global annual flux of reactive volatile organic compounds (VOCs)
355 (Guenther et al., 2006). For MICS-ASIA III, hourly biogenic emissions were provided for the
356 entire year of 2010 using the Model of Emissions of Gases and Aerosols from Nature (MEGAN)
357 version 2.04 (Guenther et al., 2006). The variables that drive MEGAN include land cover
358 information (plant function type, leaf area index) and weather condition, which includes solar

359 transmission, air temperature, humidity, wind speed, and soil moisture. In the preparation of
360 MEGAN biogenic emissions, land cover information is taken from the NASA MODIS products,
361 and weather condition are calculated using WRF simulations. Figure S1 shows biogenic
362 emissions of some selected species (isoprene and HCHO) for January 2010. High biogenic
363 emissions are found in south Asia during winter, including India, south China, and Southeast
364 Asia, where solar radiation, air temperature and vegetation covers are relatively higher than in
365 northern regions. As shown in Table 1, M1 and M5 use prescribed biogenic VOCs emissions,
366 other models except M6 use internal calculation.
367 ~~Some models used these emissions directly. Others internally calculated the biogenic emissions~~
368 ~~on line with the model predicted meteorology using the MEGAN model.~~

369 2.2.3 Biomass burning emissions

370 Biomass burning in the tropics is a strong contributor to air pollutants, and extensive biomass
371 burning in Asia, particularly Southeast Asia, exerts a great influence on air quality (Streets et al.,
372 2003). For MICS-ASIA III, biomass burning emissions were processed by re-gridding the Global
373 Fire Emissions Database version 3 (GFEDv3) (0.5 by 0.5 degree). GFED fire emissions are
374 estimated through combining satellite-detected fire activity and vegetation productivity
375 information. Carbon, dry matter, CO₂, CO, CH₄, hydrogen, nitrous oxide, NO_x, NMHC, OC, BC,
376 PM_{2.5}, total particulate matter and SO₂ emissions are estimated in monthly temporal resolution.
377 Figure S2 shows the gridded biomass burning emissions for January 2010. Biomass burning
378 activity is highest in Cambodia and some areas of Myanmar and north of Thailand (Figure S2),
379 and the peak emission season is spring. Although it has been concluded that biomass burning
380 could significantly contribute to aerosol concentrations in China, the contribution is limited for

Formatted: Highlight

381 Topic 3 study since the focused region is North China where biomass burning emissions are
382 negligible during ~~cool~~ winter (Gao et al., 2016a).

Formatted: Highlight

383 2.2.4 Volcanic SO₂ emissions

384 Volcanoes are important sources of various sulfur and halogen compounds, which play crucial
385 roles in tropospheric and stratospheric chemistry. It is estimated that SO₂ emitted from volcanoes
386 account for about 9% of the total worldwide annual SO₂ flux (Stoiber et al., 1987). The Asia-
387 Pacific region is one of the most geologically unstable regions in the world where many active
388 volcanoes are located. During MICS-ASIA Phase II, the volcano SO₂ emissions had already
389 been provided for chemical transport models (Carmichael et al. 2008b). Volcano SO₂ emissions
390 were provided, with a daily temporal resolution. In January, some volcanoes in Japan are very
391 active, such as Miyakejima (139.53°E, 34.08°N, and 775m above sea level) and Sakurajima
392 (130.65°E, 31.59°N, 1117m above sea level).

Formatted: Highlight

Formatted: Highlight

393 2.2.5 Air and Ship emissions

394 Fuel burning in aircraft and ship engines produces greenhouse gases and air pollutants. The
395 shipping and aircraft emissions used are based on HTAPv2 emission inventory (0.1 by 0.1
396 degree) for year 2010 (Janssens-Maenhout et al., 2015), provided on an annual basis. Aircraft
397 emissions include three parts: landing and takeoff (LTO), climbing and descent (CDS), and
398 cruise (CRS). Aircraft emission hot spots are mostly located in Japan, and Beijing, Yangtze
399 River Delta (YRD) and Pearl River Delta (PRD) in China (Figure S3). East China Sea, sea
400 around Japan and Singapore exhibit high shipping emissions due to active shipping
401 transportation (Figure S3). It is estimated that international shipping contributed about 10% to

Formatted: Highlight

402 the global SO₂ emissions, and together with aviation contribute more than 10% of global NO_x
403 emissions (Janssens-Maenhout et al., 2015).

Formatted: Highlight

404 2.3 Boundary conditions

405 To predict more realistic spatial and temporal variations of air pollutants, boundary conditions
406 from global chemical transport models are necessary to drive regional chemical transport models

407 (Carmichael et al., 2008b). Simulations of ~~three-two~~ global chemical transport models (i.g.,

Formatted: Highlight

408 ~~CHASER~~, GEOS-Chem and MOZART) were ~~provided-used~~ as boundary conditions for MICS-

409 ASIA III. ~~CHASER was developed in Japan to simulate the O₃-HO_x-NO_x-CH₄-CO~~

410 ~~photochemical system and its effects on climate (Sudo et al., 2002)~~. GEOS-Chem was developed

411 in the USA to simulate tropospheric chemistry driven by assimilated meteorology (Bey et al.,

Formatted: Highlight

412 2001). ~~In addition~~, the National Center for Atmospheric Research (NCAR) also provides global

413 simulations of atmospheric chemistry (MOZART model) and an interface to convert them to

414 WRF-Chem boundary conditions (Emmons et al., 2010), and NASA provides global aerosol

Formatted: Highlight

415 distributions using the global GOCART chemistry model (Chin et al., 2002). GEOS-Chem was

Formatted: Highlight

416 run with 2.5°x2° resolution and 47 vertical layers ~~and CHASER model was run with 2.8°x2.8°~~

417 ~~and 32 vertical layers. 3 hourly average fields of gaseous and aerosols were distributed to all~~

418 ~~participants~~. The MOZART-4 simulations were ~~also~~ configured at the horizontal resolution of

419 2.8°x2.8°, ~~and but~~ with 28 vertical levels. NASA GOCART was configured at the same

420 resolution as GEOS-5 meteorology (1.25°x1°). As listed in Table 1, M1 used climatological data

421 from the NOAA Aeronomy Lab Regional Oxidant Model (NALROM), while M2 used boundary

422 conditions from the MOZART-4 (provided from the NCAR website). M3 and M4 used

423 MOZART-4 as boundary conditions for gases and used GOCART as boundary conditions for

424 aerosols. M6 also used fixed climatology boundary conditions, and M5 and M7 used GEOS-

425 Chem outputs as boundary conditions. The spatial distribution of near surface concentrations of
426 major gases and aerosols from both GEOS-Chem and MOZART are shown in Figure S4. Even
427 though the same global model is used as boundary conditions, the treatments of inputs might
428 differ in details, which might lead to ~~considerable~~ dissimilarities. In MICS-ASIA II, **Holloway et**
429 **al. (2008)** discussed the impacts of uncertainties in global models on regional air quality
430 simulations.

Formatted: Highlight

431 **2.4 Observation data**

432 Historically, the lack of reliable air quality measurements in Asia has been a bottleneck in
433 understanding air quality and constraining air quality modeling in Asia. Beginning with MICS-
434 ASIA II, observational data from Acid Deposition Monitoring Network in East Asia (EANET)
435 has been used to evaluate model performance. EANET was launched in 1998 to address acid
436 deposition problems in East Asia, following the model of the Cooperative Program for
437 Monitoring and Evaluation of the Long-range Transmission of Air pollutants in Europe (EMEP).
438 As of 2010, there are 54 wet deposition sites and 46 dry deposition sites in 13 participating
439 countries. Quality assurance and quality control measures are implemented at the national levels
440 and in the Inter-laboratory Comparison Project schemes to guarantee high quality dataset.
441 EANET supported current activities of MICS-ASIA III, and provided measurements in 2010 to
442 all modeling groups. More information about EANET dataset can be found in
443 <http://www.eanet.asia/>.

444 In addition to EANET data, measurements of air pollutants and aerosol optical depth (AOD)
445 collected at the Campaign on Atmospheric Aerosol Research network of China (CARE-China)
446 (**Xin et al., 2015**) network were also used. Previous successful networks in Europe and the

Formatted: Highlight

447 United States underscored the importance of building comprehensive observational networks of
448 aerosols in China to get better understanding of the physical, chemical and optical properties of
449 atmospheric aerosols across China. As the first comprehensive attempt in China, CARE-China
450 was launched in 2011 by Chinese Academy of Sciences (CAS) (Xin et al., 2015). Before
451 launching this campaign, CAS had already been measuring air pollutants and AOD at some
452 CARE-China sites. Table 2 summaries the locations and characteristics of the CARE-China
453 measurements for January 2010. Air quality measurements include concentrations of PM_{2.5},
454 PM₁₀, SO₂, NO₂, NO, CO, O₃.

455 In addition, AOD from Aerosol Robotic Network (AERONET) (<https://aeronet.gsfc.nasa.gov/>)
456 and operational meteorological measurements (near surface temperature, humidity, wind speed
457 and downward shortwave radiation) in China and atmospheric sounding data in Beijing were
458 used. AERONET provides long-term, continuous, readily accessible and globally distributed
459 database of spectral AOD, inversion products and precipitable water. AOD data are calculated
460 for three quality levels: Level 1.0 (unscreened), Level 1.5 (cloud screened), and Level 2.0 (cloud
461 screened and quality assured) (Holben et al., 1998). The locations and characteristics of the
462 AERONET measurements are also summarized in Table 2. In-situ measurements of
463 meteorological data from standard stations in China are operated by China Meteorological
464 Administration (CMA) and different levels of data, including daily, monthly, and annually, are
465 open to the public (<http://data.cma.cn/en>). The locations of all used observational sites are
466 marked in Figure S4S5, Figure S5-S6 and Figure S6S7.

467 The meteorology measurements (locations are shown in Figure S4S5) were averaged and
468 compared with model results that averaged across those locations. The radiation measurements
469 were averaged and compared against model results in North China and South China (locations

Formatted: Highlight

470 are shown in Figure [S5S6](#)), separately. The CARE-China, AERONET and EANET
471 measurements (locations are shown in Figure [S5-S6](#) and [S6S7](#)) were compared against model
472 results site by site, and model ensemble mean values were made by averaging all model results.

473 **2.5 Analysis methodology**

474 All groups participating in Topic 3 were requested to simulate meteorology, air quality, radiative
475 forcing and effects of aerosols over the Beijing-Tianjin-Hebei region of east China during two
476 periods: January 2010 and January 2013. ~~Simulations were designed to evaluate radiative and
477 microphysical feedbacks, together and separately, relative to simulations without feedbacks.~~
478 Each group was requested to submit the following fields from their simulations.

479 (1) hourly mean meteorology:

480 (a) air temperature and water vapor mixing ratio at 2m above ground (T2, Q2), wind speed at
481 10m above ground (WS10), and shortwave radiation flux (Wm-2) at the surface;

482 (b) above variables (except shortwave radiation flux) at 1km and 3km above ground.

483 (2) hourly mean concentrations:

484 (a) SO₂, NO_x, CO, O₃, PM_{2.5}, PM₁₀ and sulfate, nitrate, ammonium, BC, OC and dust in PM_{2.5};

485 (b) above variables at 1km and 3km above ground.

486 (3) hourly mean AOD, aerosol direct radiative forcings at the surface, top of the atmosphere
487 (TOA) and inside the atmosphere (single scattering albedo is an option for participants).

488 (4) Hourly mean integrated liquid water, cloud optical depth.

489 (5) Changes in T2, Q2, WS10 and PM_{2.5} concentrations at the surface due to both direct and
490 indirect aerosol's effects.

491 We calculated multiple model evaluation metrics, including correlation coefficient (r), root mean
492 square error (RMSE), mean bias error (MBE), normalized mean bias (NMB), mean fractional
493 bias (MFB) and mean fractional error (MFE). The equations are presented in supplemental
494 information.

495

496 **3 General description of meteorology and haze during the study periods**

497 Winter haze events ~~are frequently happening~~frequently happen in east China, which is partially
498 due to the stagnant weather conditions in winter. Here we present general descriptions of the
499 meteorological conditions during ~~the selected two~~ January ~~months~~2010 using the NCEP/NCAR
500 reanalysis products. Figure ~~S8 (a, b)~~ displays the monthly mean T2 (temperature at 2m) and
501 W10 (wind speeds at 10m) ~~for January 2010 and January 2013, respectively. For both periods,~~
502 WS10 were very weak in eastern and central China regions, ~~while lower~~ T2 in Mongolia region
503 ~~was relatively higher for January 2013. Historical analyses have shown that cold conditions are~~
504 ~~usually associated with strengthened Siberian High (Gong and Ho, 2002), and relatively higher~~
505 ~~T2 and more weakened~~was associated with Siberian High. ~~As shown in Figure S8 (b), (Figure 3~~
506 ~~(e, d)) during January 2013 led to weaker winter monsoon winds and higher pollution levels. The~~
507 ~~relatively weaker Siberian High during January 2013 compared to January 2010 is also shown in~~
508 ~~the sea level pressures (Figure 3 (e, d)). The~~the Siberian High center was about ~~1037mb during~~
509 ~~January 2013, lower than that (1040mb) during January 2010. Figure 3 (e, d) show that and~~
510 there was no significant precipitation in North China and heavy rainfall only occurred in

511 Southeast Asia regions. During ~~cold~~ winters, northern China burns coal for heating, generating
512 more emissions. Under stagnant weather conditions, haze episodes are easily triggered. ~~It was~~
513 ~~reported that January 2013 was the haziest month in the past 60 years in Beijing, and~~
514 ~~instantaneous PM_{2.5} concentration exceeded 1000 μg/m³ in some areas in Beijing. Winter haze~~
515 ~~also happened from 16 to 19 January in 2010.~~ High concentrations of aerosols during ~~these two~~
516 ~~study periods~~ this month provide great opportunity to study aerosol-radiation-weather
517 interactions.

518

519 **4 Results and discussions**

520 In this section, we present some major features of model performances in meteorological and
521 chemical variables for the January 2010 period. Detail analysis of feedbacks and radiative
522 forcing are presented in MICS-ASIA III companion papers. Heavy haze occurred over broad
523 regions of East China in January 2010. The plots of observed meteorological variables and PM_{2.5}
524 in Beijing show the general situation (Figure 43). Elevated PM_{2.5} occurred during three periods
525 separated in time by roughly one week (January 8, 16 and 26). The major event occurred during
526 January 15-21. The events occurred during periods of low wind speeds, and increasing
527 temperature and relative humidity. The high PM_{2.5} concentrations during January 15-21 also
528 greatly reduce the downward shortwave radiation. Below we evaluate how well the models
529 predict these features.

530 **4.1 Evaluation of meteorological variables**

531 Air quality is affected by not only emissions, but also meteorological conditions. Meteorology
532 affects air quality through altering emissions, chemical reactions, transport, turbulent mixing,
533 and deposition processes (Gao et al., 2016b,2016c). Thus, it is important to assess how well these
534 participating models reproduced meteorological variables. The predicted temperature at 2m high
535 (T2), water vapor mixing ratio at 2m (Q2), wind speed at 10m high (WS10) and daily maximum
536 downward shortwave radiation (SWDOWN) were evaluated against near surface observations at
537 the CMA sites.

Formatted: Highlight

538 Figure 5-4 (a-c) shows the comparisons between simulated and observed daily mean T2, Q2 and
539 WS10 averaged over stations in East China (locations are shown in Figure S4S5) during January
540 2010, along with multi-model ensemble mean and observation standard deviation. The calculated
541 correlation coefficients between models and observations are also shown in Figure 5 and other
542 calculated model evaluation metrics are summarized in Table 3. In general, the simulated
543 magnitudes and temporal variations of T2 and Q2 show high order of consistencies with
544 observations, with correlation coefficients ranging from 0.88 to 1. For T2, models tend to have a
545 cool bias; M1 and M2 have the lowest RMSE (0.64 and 0.68), lowest MBE (-0.19 and -0.60) and
546 lowest NMB (-0.07% and -0.22%) values (Table 3). For Q2, most models tend to slightly
547 overestimate; M1 and M2 have the best performance, with the lowest RMSE (0.14 and 0.10),
548 lowest MBE (0.02 and -0.01), and lowest NMB (0.84% and -0.55%) values (Table 3).

549 Simulated WS10 exhibit larger diversity of results. All models tend to overestimate WS10, with
550 MBE ranging from 0.15m/s to 2.37m/s. Overestimating wind speeds under low wind conditions
551 is a common problem of current weather forecasting models, and many factors, including errors
552 in terrain data and reanalysis data, relatively low horizontal and vertical model resolutions, as
553 well as poorly parameterized urban surface effect, contribute to these overestimations. From the

554 calculated RMSE, MBE, and NMB listed in Table 3, M2, M5 and M7 show better skills in
555 capturing WS10. In addition, the multi-model ensemble mean show the lowest RMSE for Q2,
556 and also better skills than most models for T2 and WS10. The correlation coefficients between
557 multi-model ensemble mean and observations are 0.99, 0.99 and 0.98 for T2, Q2 and WS10,
558 respectively.

559 The accuracy of radiation predictions is of great significance in evaluating aerosol-radiation-
560 weather interactions. We evaluated simulated daily maximum SWDOWN averaged over sites in
561 northern China and southern China separately in January 2010 against observations. The
562 locations of the radiation sites are shown in Figure [S5S6](#). As shown in Figure [5-4](#) (d), over
563 stations in northern China, all models except M6 and M7 reproduce daily maximum SWDOWN
564 well, with correlation coefficients ranging from 0.72 to 0.94. [The poor performance of M6 in
565 North China is caused by largely overpredicted liquid water path \(LWP\) over North China
566 \(Figure S9\).](#)

567 SWDOWN decreases under conditions of high PM, as shown for example on January 9 and 15-
568 21. This is one of the important reasons for coupled air quality and meteorology modeling, as
569 they can account for this effect of aerosols. It is worth noting that most models predict higher
570 daily maximum SWDOWN compared to observations when severe haze happened in the North
571 China Plain (16-19 January 2010), indicating aerosol effects on radiation might be
572 underestimated. [Clouds are also important to alter radiation. To exclude its impacts on the
573 radiation shown here, we calculated the reduction ratio of radiation due to clouds using radiation
574 prediction for clear sky and with clouds from M2 \(shown in Figure S10\). During the severe haze
575 period \(16-19 January 2010\), the averaged reduction fraction is 5.9% in north China and 4.2% in
576 south China, suggesting the relatively lower radiation during this period shown in Figure 4\(d\) is](#)

577 mainly caused by aerosols, while the lowest radiation on 20 January was caused by clouds
578 (Figure 4(d) and Figure S10). Over southern China sites (Figure ~~5e4e~~), M6 and M7 show a better
579 consistence with observations than over northern China sites. According to the calculated RMSE
580 listed in Table 3, M3 and multi-model ensemble mean exhibit relatively better performance in
581 capturing the observed time series of daily maximum SWDOWN in both northern China and
582 southern China.

583 The above comparisons show that T2 and Q2 are reproduced well by the participating models,
584 and WS10 is overestimated by all models. Emery et al. (2001) proposed that excellent model
585 performance would be classified as wind speed RMSE smaller than 2 m/s, and wind speed bias
586 smaller than 0.5 m/s. Based on the calculated RMSE and MBE of WS10 shown in Table 3,
587 RMSE values from all models match the proposed RMSE threshold but MBE values are higher
588 than 0.5 m/s. The vertical distributions of temperature, water vapor mixing ratio and wind speeds
589 were also validated against atmospheric sounding data in Beijing at 1km and 3km (Figure
590 ~~S7S11~~, averaged at 00:00 and 12:00 UTC) (<http://weather.uwyo.edu/upperair/sounding.html>).
591 The magnitudes of temperature, water vapor mixing ratio and wind speeds from different models
592 are generally consistent with each other at 1km and 3km, but variations are larger near the
593 surface.

594 **4.2 Evaluation of air pollutants**

595 Figure ~~6-5~~ displays the daily averaged predicted and observed SO₂, NO_x, CO, O₃, PM_{2.5}, and
596 PM₁₀ at the Beijing station, along with the observation standard deviation (locations are shown in
597 Figure ~~S6S7~~). Comparisons for the Tianjin, Shijiazhuang and Xianghe sites are shown in Figure
598 ~~S8S12-S10S14~~. M6 only provided SO₂, NO_x concentrations, so it is not ~~only~~ shown in the plots

Formatted: Highlight

599 of CO, O₃, PM_{2.5}, and PM₁₀. The observed and predicted primary pollutants and PM_{2.5} and PM₁₀
600 show the same monthly variations with elevated values at roughly weekly intervals, with the
601 largest event occurring during January 15-21. For example, as shown in the comparisons of SO₂
602 concentration, the temporal variations are reproduced well by all the models, but peak values are
603 overestimated or underestimated by some models. Based on the calculated MBE values shown in
604 Table 4, all models except M2 tend to underestimate SO₂ in ~~Beijing~~ the CARE-China sites. M1
605 shows the highest correlation (0.90) with SO₂ observations in the Beijing site, and most other
606 models show similar good correlations. The multi-model ensemble mean shows a better
607 agreement with observations with a higher correlation of 0.92, and it falls within the range
608 shown with standard deviation error bar. In general, the predictions for NO_x capture the main
609 features in the observations, with slightly less skill than for the SO₂ prediction. The calculated
610 correlation coefficients for NO_x from different models are close to each other, ranging from 0.63
611 to 0.88. M2 and M5 predict higher NO_x concentrations than observations and other models
612 (MBE in Table 4). All models overestimate NO_x concentration in Shijiazhuang (Figure S8S14),
613 suggesting NO_x emissions in Shijiazhuang might be overestimated in the MIX emission
614 inventory. All models produce similar CO predictions~~All models are consistent with each in CO~~
615 ~~predictions~~.
616 PM_{2.5} concentrations are well modelled, with high correlation coefficients ranging from 0.87 to
617 0.90 in Beijing, from 0.83 to 0.93 in Tianjin, and from 0.74 to 0.91 in Xianghe. The correlation
618 coefficient of the multi-model ensemble mean for PM_{2.5} reaches 0.94 (Table 4), better than any
619 individual model. The performances of all participating models in reproducing PM₁₀ variations
620 are not as good as reproducing PM_{2.5}. M1 and M2 overestimate PM₁₀ concentrations, and other

621 models underestimate PM₁₀ concentrations (MBE in Table 4). These biases are probably related
622 to different treatments of primary aerosols and anthropogenic dust in the models.

623 The models showed the poorest skill in predicting ozone. All models exhibit different
624 performances in simulating ozone concentrations, and the correlation coefficients between
625 models and observations can reach negative values (Figure ~~S8S12~~). M3 and M4 tend to
626 overestimate ozone concentrations, M2 slightly overestimates it, and M1, M5, and M7 slightly
627 underestimate it (MBE in Table 4). According to the calculated RMSE in Table 4, M1 and M7
628 shows relatively better performance in modeling ozone variations. Although WRF-Chem and
629 NU-WRF models were applied at three institutions, different gas phase chemistry schemes were
630 used, which leads to these diversities among predicted ozone concentrations. The impacts of gas
631 phase chemical mechanisms on ozone simulations have been investigated in Zhang-Knote et al.
632 (2015), ~~but under high photochemical conditions. The results presented here winter conditions~~
633 ~~with slower photochemistry in general and where hazy conditions further reduce photochemistry~~
634 ~~through diming effects.~~

635 Figure ~~7-6~~ shows the comparisons between modeled and observed ground level daily averaged
636 concentrations of SO₂, NO_x, O₃ and PM₁₀ during January 2010 at the Rishiri site in Japan from
637 EANET. The locations of ~~used~~-EANET sites are marked in Figure ~~S6S7~~. Comparisons at other
638 EANET sites are shown in Figure ~~S11S15-S14S18~~. The models are able to predict the major
639 features in the observations. For example, low values of most pollutants are observed (and
640 predicted) during the first half of the month, followed by elevated values, which peak on January
641 21. For SO₂, most models show similar capability in producing the temporal variations in
642 observations with slight underestimation (MBE in Table 5). According to the calculated RMSE
643 averaged over all the EANET sites, M2 and the multi-model ensemble mean performed the best.
644 For NO_x, the multi-model ensemble mean shows lower RMSE than any individual model (Table
645 5). Similar to the comparisons over CARE-China sites, large discrepancies exist in ozone
646 predictions, but the model ensemble mean still shows the lowest RMSE for ozone predictions.

Formatted: Line spacing: 1.5 lines

Formatted: Font color: Blue

Formatted: Left

647 PM₁₀ concentrations are largely underestimated by M1 (largest negative MBE: -21.03ug/m³) and
648 overestimated by M5 (highest positive MBE: 3.77ug/m³) (Table 5), which could also be related
649 to differences in the way sea-salt emissions are treated in the various models. [Spatial](#)
650 [distributions of monthly near surface concentrations of SO₂, NO_x, O₃ and CO for January 2010](#)
651 [from all participating models are shown in Figure S19. The aerosol spatial distributions are](#)
652 [discussed in the following section.](#)

653

654 **4.3 PM_{2.5} and PM_{2.5} chemical composition distribution**

655 ~~Haze pollution is characterized by high loadings of PM_{2.5}, thus accurately predicting PM_{2.5} and~~
656 ~~its chemical compositions are crucial to understand haze pollution and to provide insightful~~
657 ~~implications for controlling haze in China. The accuracy of predicting PM_{2.5} chemical~~
658 ~~composition is also of great importance in estimating aerosol-radiation interactions. For example,~~
659 ~~black carbon absorbs shortwave radiation, whereas sulfate and organic carbon mostly scatter~~
660 ~~radiation.~~ Due to different implementations of chemical reactions in the models, predicted PM_{2.5}
661 chemical compositions from participating models differ largely. Figure 8-7 and Figure 9-8 show
662 the predicted monthly mean concentrations of sulfate, nitrate, ammonium, BC and OC in PM_{2.5}
663 from all participating models for January 2010.

664 M1, M2, M3, M4 and M7 all predict quite low sulfate concentrations in east China, but with
665 considerably enhanced sulfate in southwest areas of China and west areas of India. M5 and M6
666 shows similar spatial patterns of sulfate except that M6 produces higher concentrations. The
667 chemical production of sulfate is mainly from gas-phase oxidation of SO₂ by OH radicals and
668 aqueous-phase pathways in cloud water. In cloud water, dissolved SO₂ can be oxidized by O₃,

669 H₂O₂, Fe(III), Mn(II), and NO₂ (Seinfeld and Pandis, 2016). Most chemical transport models have
670 included the above gas phase oxidation of SO₂ by OH and oxidation of SO₂ by O₃ and H₂O₂ in
671 aqueous phase. Under hazy conditions, radiation is largely reduced due to aerosol dimming effects,
672 and sulfate formation from gas phase and aqueous phase oxidation processes are slowed down,
673 which tend to reduce sulfate concentration. However, field observations exhibit an increase in
674 sulfate concentration during haze episode (Zheng et al., 2015), and Cheng et al. (2016) proposed
675 that the reactive nitrogen chemistry in aerosol water could contribute significantly to the sulfate
676 increase due to enhanced sulfate production rates of NO₂ reaction pathway under high aerosol pH
677 and elevated NO₂ concentrations in the North China Plain (NCP) during haze periods. Wang et al.
678 (2016) also pointed out the aqueous oxidation of SO₂ by NO₂ is key to efficient sulfate formation
679 on fine aerosols with high relative humidity and NH₃ neutralization or under cloudy conditions.
680 Besides, Zheng et al. (2015) suggested that heterogeneous chemistry on primary aerosols could
681 play an important role in sulfate production and lead to increasing sulfate simulation during haze
682 episodes. X. Huang et al. (2014) found including natural and anthropogenic mineral aerosols can
683 enhance sulfate production through aqueous-phase oxidation of dissolved SO₂ by O₃, NO₂, H₂O₂
684 and transition metal. Gao et al. (2016), Wang et al. (2014), and Zhang et al. (2015) also emphasized
685 the importance of multiphase oxidation in winter sulfate production. However, these above
686 aqueous and heterogeneous processes are currently not incorporated in the participating models
687 for this study, which might be responsible for the apparent under-predictions of sulfate
688 concentration (Figure 499). M5 also incorporated heterogeneous chemical reactions on aerosol
689 surface (Li and Han, 2010), which enhances total sulfate production.

690 M1 and M5 predict relatively small nitrate and ammonium concentrations; while M2, M6 and
691 M7 produces similar magnitudes and spatial patterns of nitrate. Nitrate formation involves both

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

Formatted: Subscript

Formatted: Subscript

Formatted: Subscript

Formatted: Subscript

Formatted: Subscript

Formatted: Highlight

692 daytime and nighttime chemistry. During daytime, NO₂ can be oxidized by OH to form nitric
693 acid (HNO₃), and by ozone to form NO₃. HNO₃ is easily removed by dry or wet deposition, but
694 NO₃ is easily photolyzed back to NO₂. During nighttime, NO₃ is the major oxidant, which oxidizes
695 NO₂ to form dinitrogen pentoxide (N₂O₅). Homogenous reaction of N₂O₅ with water vapor is
696 possible but very slow while heterogeneous uptake of N₂O₅ onto aerosol particles has been
697 identified as a major sink of N₂O₅ and an important contributor to particulate nitrate (Kim et al.,
698 2014). The MOSAIC aerosol module (Zaveri et al., 2008) coupled with CBMZ gas phase
699 chemistry in WRF-Chem already includes heterogeneous uptake of N₂O₅ since version v3.5.1
700 (Archer-Nicholls et al., 2014), which is the version used by M2, leading to the high production of
701 nitrate. An et al. (2013) incorporated photoexcited nitrogen dioxide molecules, heterogeneous
702 reactions on aerosol surfaces, and direct nitrous acid (HONO) emissions into the WRF-Chem
703 model and found these additional HONO sources can improve simulations of HONO and nitrate
704 in north China. M7 also predict high nitrate concentrations (N₂O₅ and NO₂ gases react with
705 liquid water, Zheng et al., 2015), and the predicted lower nitrate concentrations from other
706 models are probably due to missing aqueous phase and heterogeneous chemistry, or the
707 implementations of different gas phase oxidation in these models. Many studies have been
708 conducted regarding sulfate formation issues. Nitrate also account for a large mass fraction in
709 PM_{2.5} during winter haze events in north China, yet less attention was attracted to fully
710 understand its formation. It is worth furtherly digging into the details about how different
711 processes contribute to high nitrate concentrations in future studies. M3 and M4 do not include
712 the explicit nitrate and ammonium treatment but ammonium is implicitly considered in total
713 PM_{2.5} mass estimate.

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

Formatted: Subscript

Formatted: Subscript

Formatted: Subscript

Formatted: Highlight

Formatted: Subscript

714 The predicted ammonium concentrations are associated with the amounts of sulfate and nitrate,
715 as shown by its similar spatial distribution to sulfate and nitrate. NH_3 neutralizes H_2SO_4 and
716 HNO_3 to form aerosol, so its amount can affect the formation of sulfate, nitrate and ammonium.
717 Since the same emission inventory was used, the amount of ammonia available for neutralizing
718 will not vary greatly among these models. Thus, the rates of H_2SO_4 and HNO_3 production
719 determines the amounts of ammonium. For example, the produced ammonium concentrations are
720 small in M1, similar to their sulfate and nitrate productions. High ammonium concentrations are
721 predicted from M6, due to high productions of nitrate and sulfate (Figure 87).

722 The spatial distributions and magnitudes of predicted BC from all participating models are
723 similar to each other as BC is a primary pollutant whose mass as BC is not impacted by chemical
724 reactions. The concentrations of BC in the atmosphere are mainly influenced by PBL mixing and
725 diffusion, aging, deposition (dry deposition and wet scavenging) and advection. Predicted BC
726 from M2 and M7 are higher than from other models, which might be caused by ~~different~~
727 ~~treatments of emission inventory (for example, how to distribute emissions to different vertical~~
728 ~~layers), horizontal grid interpolation, and/or different parameterizations for vertical diffusion, the~~
729 ~~treatment of aging and, deposition (dry deposition and wet scavenging) and advection processes.~~
730 For example, in the GOCART aerosol model (M3 and M4), 80% of BC are assumed to be
731 hydrophobic and then undergo aging to become hydrophilic in an e-folding time of 1.2 days.
732 Hydrophilic aerosols will go through wet deposition. But in other models like M2 and M7, BC is
733 assumed to be hydrophobic, thus the wet removal is less.

734 The disparity among predicted OC concentrations is mainly associated with the different
735 treatments of SOA production, given the POC prediction is generally consistent among models
736 using the same emission inventory. The predicted OC concentrations from M1, M2, and M7 are

737 close to each other. M1 uses SORGAM (Secondary Organic Aerosol Model) to simulate SOA, but
738 M2 and M6 did not include any SOA formation mechanism. The similar magnitudes of OC from
739 M1 ~~and M2~~ suggests that SORGAM in M1 does not produce appreciable amounts of SOA, which
740 is consistent with the findings in [Gao et al. \(2016a\)](#). Although SOA formation is implemented in
741 M5, the production is relatively weak compared to M3 and M4. In the atmosphere, SOA is mainly
742 formed from the condensation of semi-volatile VOCs from oxidation of primary VOCs. An
743 empirical 2-product model ([Odum et al., 1996](#)) is often used to simulate SOA formation, but this
744 method was reported to significantly underestimate measured SOA mass concentrations ([Heald et
745 al., 2008](#)). Later, the volatility basis-set approach ([Donahue et al., 2006](#)) was developed to
746 represent more realistically the wide range of volatility of organic compounds and more complex
747 processes, and it was found to increase SOA production and to reduce observation-simulation
748 biases in many regions with high emissions ([Tsimpidi et al., 2010](#)) including east China ([Han et
749 al., 2016](#)). It was also suggested that primary organic aerosols (POA) are semi-volatile and can
750 evaporate to become SOA precursors, which promotes the understanding and improvements of
751 SOA modeling ([Li et al., 2014](#); [Kanakidou et al., 2005](#)). In M5, the SOA production is calculated
752 using a bulk yield method via [Lack et al. \(2004\)](#), in which the amount of SOA able to be produced
753 from a unit of reacted VOC from anthropogenic and biogenic origins are used to represent SOA
754 yields. However, the SOA concentration is highly dependent on the yield data. During haze
755 episodes, photochemistry is reduced due to the aerosol dimming effect, thus aqueous reaction
756 processes on aerosol water and cloud/fog water could become much more important in producing
757 SOA, ~~as suggested in Cheng et al. (2016)~~. [R. Huang et al. \(2014\)](#) also suggested that low
758 temperature does not significantly reduce SOA formation rates of biomass burning emissions.
759 Most models over-simplified SOA formation.

Formatted: Highlight

760 ~~The missing representation of such process in the participating models may partly account for the~~
761 ~~low values in the simulated SOA.~~ In M3 and M4, SOA is treated by assuming that 10% of VOCs
762 from terrestrial source are converted to OC (Chin et al., 2002), and these models produce high OC
763 concentrations, with a major contribution from SOA. The 10% yield rate could be unrealistically
764 high during hazy days because solar radiation was much reduced. Zhao et al. (2015)
765 comprehensively assessed the effect of organic aerosol aging and intermediate-volatile emissions
766 on OA formation and confirmed their significant roles. All these results suggest more complicated
767 SOA scheme are needed to improve organic aerosol simulations during haze events.

Formatted: Highlight

768 The different predictions of PM_{2.5} chemical components lead to differences in PM_{2.5} ~~and~~ PM₁₀
769 concentrations for January 2010, which are shown in the last row of Figure 98. Although spatial
770 distributions of PM_{2.5} from these models are similar, the underlying causes are different. M2, M3
771 and M5 simulated higher PM_{2.5} levels in deserts of west China, which are contributed by wind-
772 blown dust deflation. M1 and M7 fail to produce high PM_{2.5} concentrations in the deserts of west
773 China, due to omission of dust emissions. M4 presented results in a smaller domain excluding
774 west China. The enhanced PM_{2.5} concentrations in Central China from M2 and M7 are caused by
775 large nitrate production, as shown in Figure 87.

Formatted: Indent: Left: 0 cm, Hanging: 0.5 ch, First line: -0.5 ch

Formatted: Subscript

776 The differences in the predictions of aerosols composition discussed above can be seen clearly in
777 the comparisons at the Beijing site on 13-23 January when a haze event occurred in the NCP
778 (Figure 499). Also shown are the observed values. Most models fail to produce the observed
779 high sulfate concentrations. Only the sulfate predictions from M5 are close to the observed
780 values. Sulfate is much lower than observed for all other models, except M6 which is too high.
781 M2 and M7 predict reasonable nitrate concentrations. M3 and M4 overpredict OC during the
782 haze period, but other models underpredict OC concentrations. Only M5 prediction is close to

783 ~~observation, and M6 predicts higher sulfate level. M2 and M7 predict reasonable nitrate~~
784 ~~concentrations. M3 and M4 largely overpredict OC during haze period, but other models tend to~~
785 ~~underpredict OC concentrations.~~

786 Figure ~~11-10~~ and ~~112~~ show the ensemble mean monthly averaged near-surface PM_{2.5}, PM_{2.5}
787 composition, along with the spatial distribution of the coefficient of variation. The coefficient of
788 variation is defined as the standard deviation divided by the average (Carmichael et al., 2008b),
789 and larger values indicate lower consistency among models. Mean concentrations of PM_{2.5} and
790 PM_{2.5} chemical compositions are high in Sichuan Basin and east China. High coefficient of
791 variation are shown in North China for sulfate, and in most areas for nitrate and OC. The diversity
792 in predictions of these species are caused by complexity of secondary formation and different
793 treatments in models as discussed earlier. Higher consistency is shown for model BC with
794 coefficient of variations less than 0.3 in most areas (Figure 10-(h)). Coefficient of variations for
795 PM_{2.5} are also low in North China region, which is consistent with good performance of PM_{2.5}
796 predictions shown in above comparisons. However, the coefficient of variation can reach above
797 1.6 in northwestern regions, partially due to discrepancies in dust predictions.

798 4.4 Evaluation of AOD

799 AOD is an indication of aerosol pollution, which tells us how much sunlight is blocked from
800 reaching the surface by suspended aerosols. We used the measurements of AOD at AERONET
801 and CARE-China sites to evaluate how participating models perform in simulating AOD. ~~In~~
802 ~~WRF-Chem, AOD is usually calculated at 300, 400, 600 and 999nm, which can be converted to~~
803 ~~AOD at other wavelengths based on Angstrom exponent relation (Schuster et al., 2006).~~ The
804 submitted AOD from all models except M6 are at 550nm, and AOD from M6 are at 495nm. We

Formatted: Highlight

Formatted: Highlight

805 used Angstrom exponent relation (Schuster et al., 2006) to convert AOD from M6 at 495nm to
806 550nm, and all used AERONET and CARE-China AOD data to 550nm. The locations of
807 AERONET and CARE-China AOD measurement sites are shown in Figure S5. Daytime mean
808 AOD are calculated in pairwise manner and the comparisons and performance statistics are
809 shown in Figure 4312, 4413, and Table 6. On some days, data are missing because AOD cannot
810 be retrieved under serious pollution and cloudy conditions (Gao et al., 2016a). On days with
811 data, the variations of AOD are captured well by all models. However, large disparities exist
812 among models in the simulated peak AOD values (factor of 2) at monitoring stations during the
813 severe haze episode on 15-20 January 2010 (Figure 43-12 and Figure 4413). The participating
814 models exhibit various skill in simulating AOD temporal variation at different sites.

Formatted: Highlight

815 At CARE-China sites, M7 produces the best correlation coefficient R (0.83) among models at
816 Baoding and Beijing forest sites, M2 produces the highest R (0.86) at Cangzhou site, whereas
817 M5 shows the highest R (0.93) at the Beijing city site. At AERONET sites, M7 shows the
818 highest R (0.81) at Beijing, whereas M2 and M5 produce R as high as 0.91 at Xianghe site,
819 which is about 60km southeast of downtown Beijing. In terms of AOD magnitude, it's it is
820 interesting to note that during the severest haze days around 19 January 2010, M2 consistently
821 simulates the highest AOD among models, followed by M5 and M7, with the lowest AOD
822 from M6 simulates the lowest, and other models in the middle at the sites (Baoding, Beijing City,
823 Beijing Forest, Cangzhou, Beijing, Xianghe) in the north China plain (NCP). It is important to
824 explore the causes for the disparities in AOD predicitions.

Formatted: Highlight

~~825 AOD is calculated as the vertical integration of extinction coefficient, which is a function of
826 particle mass extinction efficiency (extinction cross section) and mass concentration. The
827 extinction efficiency is determined by particle size, refractive index and wave length. Aerosol size~~

828 ~~can grow bigger as ambient relative humidity increases, which is known as aerosol hygroscopic~~
829 ~~growth. The overall extinction coefficient of all aerosols also depends on mixing state among~~
830 ~~aerosols. Therefore, AOD is closely related a series of affecting factors from both aerosol physical~~
831 ~~properties, mass concentration and meteorological conditions.~~

832 In M1, M5, M6 and M7, particle size distribution is described by a lognormal function with a
833 geometric mean radius and a geometric standard deviation basically based on OPAC (Optical
834 properties of aerosols and clouds) database (Hess et al. 1998). In M3 and M4, sulfate, BC and OC
835 are parameterized in bulk mode, and a sectional scheme is used for sea-salt and dust aerosols. M2
836 uses an 8 bins sectional aerosol scheme with size sections ranging from 39nm to 10 μ m. The
837 refractive index of various aerosol components in the models are mainly taken from d'Almeida et
838 al. (1991) or OPAC database. All models except M6 use a kappa (κ) parameterization (Petters and
839 Kreidenweis, 2007), in which the aerosol hygroscopicity κ largely varies among different aerosol
840 chemical components, such as $\kappa=0$ for black carbon, and $\kappa>0.6$ for inorganic aerosols, but the
841 prescribed κ values could be different in the above models. M6 uses a different hygroscopic growth
842 scheme following Kiehl and Briegleb (1993). WRF-Chem models assume internally mixing
843 among aerosols within each mode (or size bin) and externally mixing between modes (or size bins),
844 M5 assumes inorganic and carbonaceous aerosols are internally mixed and externally mixed with
845 soil dust and sea-salt. M6 uses an external mixture assumption among aerosols except for
846 hydrophilic BC, which is internally mixed with other aerosols in a core-shell way.

847 We first look at the mass concentrations of different aerosol components because of their important
848 roles in determining optical properties. The observed total inorganic aerosol concentration in
849 Beijing on 19 January 2010 is about 130 μ g/m³, with sulfate and nitrate being about 50 and 65 μ g/m³,

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

Formatted: Highlight

850 respectively (Figure 409). The models generally predict a much lower sulfate concentration except
851 that the prediction from M5, which is close to observations, and M6, which shows an
852 overprediction. Most models predict lower nitrate concentration, in contrast to the overprediction
853 by M2. In terms of inorganic aerosols, which have a similar optical properties, the total
854 concentration (the sum of sulfate, nitrate and ammonium) from M2 ($175\mu\text{g}/\text{m}^3$) is higher than
855 observation and other models, and this can explain the largest simulated AOD by M2. M6
856 simulates a similar level of inorganic aerosols to M2, but the simulated AOD is lower than other
857 models, which could be due to a weak hygroscopicity or lower simulated RH-simulation (see
858 Figure S14S20). For example, high RH on January 19 are captured by M2 and M6, but
859 underpredicted by M6 (Figure S14a). Although M3 and M4 largely overpredict OC concentrations,
860 their simulated AOD are lower than M1 and M5 because their simulated inorganic aerosol
861 concentrations are much lower and OC has a smaller (mass) extinction coefficient than inorganic
862 aerosols. M1 predicts about three times larger BC concentration than the observations, although
863 the mass extinction coefficient of BC is even larger than inorganic aerosols, the mass concentration
864 and hygroscopicity of BC are much smaller and weaker than that of inorganic aerosols, leading to
865 relatively lower AOD from M1 simulation. M5 and M7 predict a similar level of inorganic aerosol
866 concentrations ($80\sim 90\mu\text{g}/\text{m}^3$) and use a similar hygroscopic growth scheme, and this can help
867 explain their consistency in the simulated AOD magnitude.

868 As listed in Table 1, only M6 uses external mixing for aerosols, and internal mixing is assumed by
869 other models for major aerosol compositions. Curci et al. (2015) discussed the impacts of mixing
870 state on simulated AOD and found that external mixing state assumption significantly increase
871 simulated AOD. M6 used external mixing but shows a relative lower AOD because of ignorance
872 of other aerosol species like dust, sea-salt, etc. In general, it appears the magnitude of inorganic

Formatted: Highlight

873 aerosol concentrations and the hygroscopic growth efficiency (affected by varied RH) can account
874 for or explain the simulated variations and magnitudes of AOD in Beijing during the severe haze
875 event, given the aerosol size ~~distribution-lognormal treatments~~ and mixing state are alike among
876 models.

877 Table 6 shows the statistics for AOD simulation at NCP sites and at all sites. In the NCP region,
878 R ranges from 0.36~0.74 for all the models. M2, M5 and M7 produce R around 0.7, indicating a
879 better simulation of AOD variations. M2 and M7 exhibit the best R (0.65) for all sites. It ~~is's~~
880 noteworthy that R values at the sites in NCP are larger than that at all sites, indicating the larger
881 reliability of model inputs (emissions and boundary conditions) and meteorological simulations.
882 In terms of magnitudes, all models tend to underpredict AOD in the whole domain, with NMB of
883 -2.7 to -71% in the NCP, and larger biases (NMB of -21~-75%) at all sites. M7 shows the smallest
884 MBE (-0.05) and NMB (-2.7%) and M2 produces the smallest RMSE. It is interesting to note that
885 the simulated AOD from the WRF-Chem models differed largely (-12 to -71%) between M1 and
886 M3 at the NCP sites, and the WRF-Chem model using finer grid size (M4) can produced slightly
887 smaller NMB compared with the same model using larger grid size (M3). However, as grid size
888 becomes finer, R and RMSE from M4 may become worse, although AOD magnitude improved.
889 The effect of grid resolution will be a topic of future paper.

890

891 **5 Summary**

892 The MICS-Asia Phase III Topic 3 examines how current online coupled air quality models
893 perform in reproducing extreme aerosol pollution episodes in North China, and how high aerosol
894 loadings during these episodes interact with radiation and weather. ~~Two hazy winter months,~~

895 ~~namely January 2010 and January 2013, were studied by six modeling groups from China, Korea~~
896 ~~and the United States.~~ Predicted meteorological variables and air pollutants from these modeling
897 groups were compared against each other, and measurements as well. A new anthropogenic
898 emission inventory was developed for this phase (Li et al., 2017), and this inventory along with
899 biogenic, biomass burning, air and ship, and volcano emissions were provided to all modeling
900 groups. All modelling groups were required to submit results based on the analysis methodology
901 that documented in this paper. This paper focused on the evaluation of the predictions of
902 meteorological parameters and the predictions of aerosol mass, composition and optical depth.
903 These factors play important roles in feedbacks impacting weather and climate through radiative
904 and microphysical processes.

905 Comparisons against daily meteorological variables demonstrated that all models can capture the
906 observed near surface temperature and water vapor mixing ratio, but near surface wind speeds
907 ~~are were~~ overestimated by all models to varying degrees. The observed daily maximum
908 downward shortwave radiation, particularly low values during haze days, were represented in the
909 participating models. Comparisons with measurements of air pollutants, including SO₂, NO_x,
910 CO, O₃, PM_{2.5}, and PM₁₀, from CARE-China and EANET networks showed that the main
911 features of accumulations of air pollutants are represented in current generation of online
912 coupled air quality models. The variations in observed AOD from CARE-China and AERONET
913 networks were also reproduced by the participating models. Differences ~~were found exist~~
914 between simulated air pollutants, particularly ozone. While winter time ozone levels are
915 typically low (below 40 ppb) as photochemical pathways are slow, the models captured the
916 synoptic variability but differed in the absolute magnitudes of near surface concentrations. The
917 role of dry deposition and the boundary conditions play important roles, which are probably

Formatted: Highlight

918 ~~related to different treatments of emission inventory, different meteorological and chemical~~
919 ~~parameterizations, and uncertainties in interpolations from original emission inventory to model~~
920 ~~grids might also contribute to these differences.~~

921 ~~Manifol~~Large differences in the models~~ld diversities~~ were found in the predicted PM_{2.5} chemical
922 compositions, especially secondary inorganic ~~aerosols~~ and organic carbon. During winter haze
923 events, the production from gas phase chemistry is inhibited, and whether including other aerosol
924 formation pathways (such as aqueous phase chemistry), or how these chemistry is parametrized
925 leads to the large difference between simulated concentrations of secondary inorganic aerosols.
926 In addition, differences in treatments of SOA also lead to large discrepancies between simulated
927 OC concentrations. Differences in the simulated variations and magnitudes of AOD in Beijing
928 during the severe haze event could be explained by the differences in simulated inorganic aerosol
929 concentrations and the hygroscopic growth efficiency (affected by varied RH).

930 Results of this intercomparison show that there remain important issues with current coupled
931 models in predicting winter haze episodes. Low wind speeds play an important role in haze
932 episodes. Current models can predict the low wind speed - high haze relationship, but
933 overestimate the low wind speeds. This contributed to the underestimation of PM_{2.5}. The models
934 also underestimate the production of secondary inorganic aerosols. There is currently a great deal
935 of research focused on inorganic aerosol production under winter haze conditions and new
936 pathways need to be included in the models to improve prediction skills. Furthermore, current
937 models have various treatments of SOA production and these lead to wide differences in the
938 contribution of SOA to winter haze episodes.

Formatted: Subscript

939 However, it was also found that using the ensemble mean of the models produced the best
940 prediction skill. While this has been shown for other conditions (for example prediction of high
941 ozone events in the US (McKeen et al., 2004), this is to our knowledge the first time it has been
942 shown for heavy haze events.

Formatted: Highlight

943 The uncertainties in predictions in aerosols mass and composition will impact estimates of the
944 aerosol direct and indirect effects during haze event (Gao et al., 2017a, 2017b, 2017c). The
945 results of the MICS-Asia Topic 3 experiments looking at the direct and indirect effects during
946 these heavy haze events is the subject of companion papers.

Formatted: Highlight

947 ~~These results provide some directions for future model improvements, and underscore the~~
948 ~~importance of accurately predicting aerosol concentration and compositions. Differences in the~~
949 ~~simulated variations and magnitudes of AOD in Beijing during the severe haze event could be~~
950 ~~explained by the differences in simulated inorganic aerosol concentrations and the hygroscopic~~
951 ~~growth efficiency (affected by varied RH).~~

952 ~~Previous studies have studied radiative forcing during haze event (Gao et al., 2017), but there are~~
953 ~~large uncertainties in aerosol modeling during haze events and in estimating its interactions with~~
954 ~~weather and climate. The uncertainties come from model inputs (land use data, model initial and~~
955 ~~boundary conditions, etc.), physical and chemical mechanisms, and particularly~~
956 ~~parameterizations of aerosol radiation weather interactions. Other companion papers with~~
957 ~~respect to MICS-Asia Topic 3 will provide more insights into current knowledge of aerosol-~~
958 ~~weather interactions under heavy pollution conditions.~~

959

960 **ACKNOWLEDGMENTS**

961 The authors would like to acknowledge support of this project from National Natural Science
962 Foundation of China ([NSFC \(41620104008\) No. 91644217](#)), and ground measurements from
963 Yuesi Wang's research group. The ground observation was supported by the National Natural
964 Science Foundation of China (41222033; 41375036) and the CAS Strategic Priority Research
965 Program Grant (XDA05100102, XDB05020103).

966

967

968

969

970

971

972

973

974

975

976

977

978

979 **Reference**

980 Akimoto, H. (2003). "Global Air Quality and Pollution." Science, 302 (5651), 1716-1719.

Formatted: Font: 小四

Formatted: Font: 小四

981 An, J., Li, Y., Chen, Y., Li, J., Qu, Y. and Tang, Y., 2013. Enhancements of major aerosol
982 components due to additional HONO sources in the North China Plain and implications for
983 visibility and haze. Advances in Atmospheric Sciences, 30(1), p.57.

Formatted: Automatically adjust right indent when grid is defined, Space After: 10 pt, Adjust space between Latin and Asian text, Adjust space between Asian text and numbers

Formatted: Font: 小四, No underline

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

984 Archer-Nicholls, S., et al. (2014). "Gaseous chemistry and aerosol mechanism developments for
985 version 3.5.1 of the online regional model, WRF-Chem." Geoscientific Model Development7(6):
986 2557-2579.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

987 Baklanov, A., et al. (2014). "Online coupled regional meteorology chemistry models in Europe:
988 current status and prospects." Atmospheric Chemistry and Physics14(1): 317-398.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

989 Baklanov, A., Brunner, D., Carmichael, G., Flemming, J., Freitas, S., Gauss, M., Hov, Ø.,
990 Mathur, R., Schlünzen, K.H., Seigneur, C. and Vogel, B., 2017. Key issues for seamless
991 integrated chemistry-meteorology modeling. Bulletin of the American Meteorological Society,
992 (2017).

993 Bey, I., et al. (2001). "Global modeling of tropospheric chemistry with assimilated meteorology:
994 Model description and evaluation." Journal of Geophysical Research: Atmospheres106(D19):
995 23073-23095.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

996 Carmichael, G. P., L.R. (1984). "An Eulerian transport/transformation/removal model for SO₂
997 and sulfate-I. model development." Atmospheric Environment18(5): 937-951.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

998 Carmichael, G. P., L.R.; Kitada, T. (1986). "A second generation model for regional-scale
999 transport/chemistry/deposition." Atmospheric Environment20(1): 173-188.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1000 Carmichael, G. R., Peters, L.R.; Saylor, R. D. (1991). "The STEM-II regional scale acid
1001 deposition and photochemical oxidant model-I. an overview of model development and
1002 applications." Atmospheric Environment25A(10): 2077-2090.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1003 Carmichael, G. R., et al. (1998). "Tropospheric ozone production and transport in the springtime
1004 in east Asia." Journal of Geophysical Research: Atmospheres103(D9): 10649-10671.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1005 Carmichael, G. R. C., G.; Hayami, H.; Uno, I.; Cho, S.Y.; Engardt, M.; Kim, S.B.; Ichikawa, Y.;
1006 Ikeda, Y.; Woo, J.H.; Ueda, H.; Amann, M. (2002). "The MICS-Asia study: model
1007 intercomparison of long-range transport and sulfur deposition in East Asia." Atmospheric
1008 Environment36: 175-199.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1009 Carmichael, G. R., et al. (2008a). "Predicting air quality: Improvements through advanced
1010 methods to integrate models and measurements." Journal of Computational Physics227(7): 3540-
1011 3571.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1012 Carmichael, G., et al. (2008b). "MICS-Asia II: The model intercomparison study for Asia Phase
1013 II methodology and overview of findings." Atmospheric Environment42(15): 3468-3490.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1014 Carter, W.P., 2000a. Documentation of the SAPRC-99 chemical mechanism for VOC reactivity
1015 assessment. Contract, 92(329), pp.95-308.

1016 Carter, W.P., 2000b. Implementation of the SAPRC-99 chemical mechanism into the models-3
1017 framework. Report to the United States Environmental Protection Agency, January, 29.

Formatted: Indent: Left: 0 cm, Automatically adjust right indent when grid is defined, Space After: 10 pt, Adjust space between Latin and Asian text, Adjust space between Asian text and numbers

1018 Cheng, Y. F., Z., G.; Wei, C.; Mu, Q.; Zheng, B.; Wang, Z.; Gao, M.; Zhang, Q.; He, K.;
1019 Carmichael, G.; Poschl, U.; Su, Hang (2016). "Reactive nitrogen chemistry in aerosol water as a
1020 source of sulfate during haze events in China." Science Advances 2(e1601530)
1021 Chin, M. G., P.; Kinne, S.; Torres, O.; Holben, B.N.; Duncan, B.N.; Martin, R.V.; Logan, J.A.;
1022 Higurashi, A.; Nakajima, T. (2002). "Tropospheric aerosol optical thickness from the GOCART
1023 Model and Comparisons with satellite and sun photometer measurements." Journal of
1024 Atmospheric Sciences(59).
1025 Curci, G., Hogrefe, C., Bianconi, R., Im, U., Balzarini, A., Baró, R., Brunner, D., Forkel, R.,
1026 Giordano, L., Hirtl, M. and Honzak, L., 2015. Uncertainties of simulated aerosol optical
1027 properties induced by assumptions on aerosol physical and chemical properties: An AQMEII-2
1028 perspective. Atmospheric Environment, 115, pp.541-552,
1029 Donahue, N.M., Robinson, A.L., Stanier, C.O. and Pandis, S.N., 2006. Coupled partitioning,
1030 dilution, and chemical aging of semivolatile organics. Environmental Science &
1031 Technology, 40(8), pp.2635-2643,
1032 D’Almeida, G. A., P. Koepke, and E. P. Shettle (1991), Atmospheric Aero-sols: Global
1033 Climatology and Radiative Characteristics,A.Deepak, Hampton, Va.
1034 _____
1035 Emery, C. T., E.; Yarwood, G. (2001). Enhanced meteorological modeling and performance
1036 evaluation for two Texas ozone episodes.
1037 Emmons, L. K. W., S.; Hess, P.G.; Lamarque, J-F.; Pfister G.G.; Fillmore, D.; Granier, C.;
1038 Guenther, A.; Kinnison, D.; Laeple, T.; Orlando, J.; Tie, X.; Tyndall, G.; Wiedinmyer, C.;

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Indent: Left: 0 cm, First line: 0 cm, Automatically adjust right indent when grid is defined, Space After: 10 pt, Adjust space between Latin and Asian text, Adjust space between Asian text and numbers

Formatted: Indent: Left: 0 cm, Automatically adjust right indent when grid is defined, Space After: 10 pt, Adjust space between Latin and Asian text, Adjust space between Asian text and numbers

1039 Baughcum, S.L.; Kloster, S. (2010). "Description and evaluation of the Model for Ozone and
 1040 Related chemical Tracers, version 4 (MOZART-4)." *Geoscientific Model Development*3: 43-67.

1041 Fast, J.D., Gustafson, W.I., Easter, R.C., Zaveri, R.A., Barnard, J.C., Chapman, E.G., Grell, G.A.,
 1042 and Peckham, S.E., 2006. Evolution of ozone, particulates, and aerosol direct radiative forcing in
 1043 the vicinity of Houston using a fully coupled meteorology - chemistry - aerosol model. *Journal*
 1044 of *Geophysical Research: Atmospheres*, 111(D21).

1045 Forkel, R., Balzarini, A., Baró, R., Bianconi, R., Curci, G., Jiménez-Guerrero, P., Hirtl, M.,
 1046 Honzak, L., Lorenz, C., Im, U. and Pérez, J.L., 2015. Analysis of the WRF-Chem contributions
 1047 to AQMEII phase2 with respect to aerosol radiative feedbacks on meteorology and pollutant
 1048 distributions. *Atmospheric Environment*, 115, pp.630-645.

1049 Galmarini, S., Koffi, B., Solazzo, E., Keating, T., Hogrefe, C., Schulz, M., Benedictow, A.,
 1050 Griesfeller, J.J., Janssens-Maenhout, G., Carmichael, G. and Fu, J., 2017. Coordination and
 1051 harmonization of the multi-scale, multi-model activities HTAP2, AQMEII3, and MICS-Asia3:
 1052 simulations, emission inventories, boundary conditions, and model output formats. *Atmospheric*
 1053 *Chemistry and Physics*, 17(2), pp.1543-1555.

1054 Gao, M., et al. (2015). "Health impacts and economic losses assessment of the 2013 severe haze
 1055 event in Beijing area." *Sci Total Environ*511: 553-561.

1056 Gao, M., et al. (2016a). "Modeling study of the 2010 regional haze event in the North China
 1057 Plain." *Atmospheric Chemistry and Physics*16(3): 1673-1691.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1058 [Gao, M., et al. \(2016b\). "Improving simulations of sulfate aerosols during winter haze over](#)
1059 [Northern China: the impacts of heterogeneous oxidation by NO2." Frontiers of Environmental](#)
1060 [Science & Engineering](#)10(5).

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1061 [Gao, M., et al. \(2016c\). "Response of winter fine particulate matter concentrations to emission](#)
1062 [and meteorology changes in North China." Atmospheric Chemistry and Physics](#)16(18): 11837-
1063 [11851.](#)

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1064 [Gao, M., et al. \(2017a\). "Chemical and Meteorological Feedbacks in the Formation of Intense](#)
1065 [Haze Events." Air Pollution in Eastern Asia: An Integrated Perspective. Springer, Cham.](#) 437-
1066 [452.](#)

1067 [Gao, M., et al. \(2017b\). Distinguishing the roles of meteorology, emission control measures,](#)
1068 [regional transport, and co-benefits of reduced aerosol feedbacks in "APEC Blue". Atmospheric](#)
1069 [Environment.](#) 167, 476-486.

1070 [Gao, M., et al. \(2017c\). "Estimates of Health Impacts and Radiative Forcing in Winter Haze in](#)
1071 [eastern China through constraints of surface PM2.5 predictions." Environ Sci Technol.](#)

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1072 [Gery, M. W. W., G.Z.; Killus, J.P.; Dodge, M.C. \(1989\). "A photochemical kinetics mechanism](#)
1073 [for urban and regional scale computer modeling " Journal of Geophysical Research](#)94(D10):
1074 [12925-12956.](#)

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1075 [Grell, G. A., et al. \(2005\). "Fully coupled "online" chemistry within the WRF model."](#)
1076 [Atmospheric Environment](#)39(37): 6957-6975.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1077 Guenther, A. K., T.; Harley, P.; Wiedinmyer, C.; Palmer, P.I.; Geron, C. (2006). "Estimates of
1078 global terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols
1079 from Nature)." Atmospheric Chemistry and Physics6: 3181-3210.

1080 Han, Zhiwei, Hiromasa Ueda, Kazuhide Matsuda, Renjian Zhang, Kimio Arao, Yutaka Kanai,
1081 Hisashi Hasome. 2004. Model study on particle size segregation and deposition during Asian
1082 dust events in March 2002, Journal of Geophysical Research, 109, D19205, doi:
1083 10.1029/2004jd004920.

1084 Han, Zhiwei. (2010). "Direct radiative effect of aerosols over East Asia with a Regional coupled
1085 Climate/Chemistry model." Meteorologische Zeitschrift, 19(3): 287-298.

1086 Han Zhiwei, Jiawei Li, Xiangao Xia, Renjian Zhang, 2012. Investigation of direct radiative
1087 effects of aerosols in dust storm season over East Asia with an online coupled
1088 regional climate-chemistry-aerosol model. Atmospheric Environment, 54, 688-699.

1089 Han Zhiwei, Jiawei Li, Weidong Guo, Zhe Xiong, Wu Zhang, 2013. A study of dust radiative
1090 feedback on dust cycle and meteorology over East Asia by a coupled regional climate-chemistry-
1091 aerosol model. Atmospheric Environment, 68, 54-63.

1092 Han Zhiwei et al.,2016. Modeling organic aerosols over east China using a volatility basis-set
1093 approach with aging mechanism in a regional air quality model. Atmospheric Environment 124,
1094 186-198.

1095 Heald, C.L., Henze, D.K., Horowitz, L.W., Feddema, J., Lamarque, J.F., Guenther, A., Hess,
1096 P.G., Vitt, F., Seinfeld, J.H., Goldstein, A.H. and Fung, I., 2008. Predicted change in global

Formatted: Default Paragraph Font, Font: (Default) +Body (Calibri), 小四, Pattern: Clear

Formatted: Font: 小四, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Default Paragraph Font, Font: (Default) +Body (Calibri), 小四, Pattern: Clear

Formatted: Font: 小四, Pattern: Clear

Formatted: Default Paragraph Font, Font: (Default) +Body (Calibri), 小四, Pattern: Clear

Formatted: Font: 小四, Pattern: Clear

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

1097 secondary organic aerosol concentrations in response to future climate, emissions, and land use
1098 change. Journal of Geophysical Research: Atmospheres, 113(D5).

1099 Hess, M., Koepke, P., Schuit, I., 1998. Optical properties of aerosols and clouds: the software
1100 package OPAC. Bull. Am. Meteorol. Soc. 79, 831-844.

1101 Holben, B. N., Eck, T.F., Slutsker, I., Tanre, D., Buis, J.P., Setzer, A., Vermote, E., Reagan, J.A.,
1102 Kaufman, Y.J., Nakajima, T. and Lavenu, F. (1998). "AERONET—A federated instrument
1103 network and data archive for aerosol characterization." Remote sensing of environment66(1): 1-
1104 16.

1105 Holloway, T., et al. (2008). "MICS-Asia II: Impact of global emissions on regional air quality in
1106 Asia." Atmospheric Environment42(15): 3543-3561.

1107 Huang, M., Carmichael, G.R., Pierce, R.B., Jo, D.S., Park, R.J., Flemming, J., Emmons, L.K.,
1108 Bowman, K.W., Henze, D.K., Davila, Y. and Sudo, K., 2017. Impact of intercontinental
1109 pollution transport on North American ozone air pollution: an HTAP phase 2 multi-model
1110 study. Atmospheric Chemistry and Physics, 17(9), pp.5721-5750.

1111 Huang, X., Song, Y., Zhao, C., Li, M., Zhu, T., Zhang, Q. and Zhang, X., 2014. Pathways of
1112 sulfate enhancement by natural and anthropogenic mineral aerosols in China. Journal of
1113 Geophysical Research: Atmospheres, 119(24).

1114 Huang, R.J., Zhang, Y., Bozzetti, C., Ho, K.F., Cao, J.J., Han, Y., Daellenbach, K.R., Slowik,
1115 J.G., Platt, S.M., Canonaco, F. and Zotter, P., 2014. High secondary aerosol contribution to
1116 particulate pollution during haze events in China. Nature, 514(7521), pp.218-222.

Formatted: Font: 小四, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Pattern: Clear

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Indent: Left: 0 cm

1117 Jacobson, M. Z., 2001: Global direct radiative forcing due to multicomponent anthropogenic and
1118 natural aerosols. J. Geophys. Res., 106, 1551–1568.

Formatted: Font: 小四

1119 Janssens-Maenhout, G., et al. (2015). "HTAP v2.2: a mosaic of regional and global emission
1120 grid maps for 2008 and 2010 to study hemispheric transport of air pollution." Atmospheric
1121 Chemistry and Physics15(19): 11411-11432.

Formatted: Font: 小四

1122 Kanakidou, M., Seinfeld, J.H., Pandis, S.N., Barnes, I., Dentener, F.J., Facchini, M.C.,
1123 Dingenen, R.V., Ervens, B., Nenes, A.N.C.J.S.E., Nielsen, C.J. and Swietlicki, E., 2005. Organic
1124 aerosol and global climate modelling: a review. Atmospheric Chemistry and Physics, 5(4),
1125 pp.1053-1123.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

1126 Kim, S. W., et al. (2009). "NO₂ columns in the western United States observed from space and
1127 simulated by a regional chemistry model and their implications for NO_x emissions." Journal of
1128 Geophysical Research114(D11).

Formatted: Font: 小四

1129 Kim, Y. J. S., S.N.; Carmichael, G.R.; Riemer, N.; Stanier, C.O. (2014). "Modeled aerosol
1130 nitrate formation pathways during wintertime in the Great Lakes region of North America."
1131 Journal of Geophysical Research: Atmospheres119: 12420-12445.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1132 Kiehl, J.T., Briegleb, B.P., 1993. The relative roles of sulfate aerosols and greenhouse gases in
1133 climate forcing. Science 260, 311-314.

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

1134 Knote, C., et al. (2015). "Influence of the choice of gas-phase mechanism on predictions of key
1135 gaseous pollutants during the AQMEII phase-2 intercomparison." Atmospheric Environment115:
1136 553-568.

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

1137 Lack, D. A., et al. (2004). "Seasonal variability of secondary organic aerosol: A global modeling
 1138 study." Journal of Geophysical Research: Atmospheres, 109(D3): n/a-n/a.

1139 Lelieveld, J., et al. (2015). "The contribution of outdoor air pollution sources to premature
 1140 mortality on a global scale." Nature, 525(7569): 367-371.

1141 Li, M., et al. (2017). "MIX: a mosaic Asian anthropogenic emission inventory under the
 1142 international collaboration framework of the MICS-Asia and HTAP." Atmospheric Chemistry
 1143 and Physics, 17(2): 935-963.

1144 Li, J. and Han, Z., 2016. Aerosol vertical distribution over east China from RIEMS-Chem
 1145 simulation in comparison with CALIPSO measurements. Atmospheric Environment, 143,
 1146 pp.177-189.

1147 McKeen, S., Wilczak, J., Grell, G., Djalalova, I., Peckham, S., Hsie, E.Y., Gong, W., Bouchet,
 1148 V., Menard, S., Moffet, R. and McHenry, J., 2005. Assessment of an ensemble of seven real -
 1149 time ozone forecasts over eastern North America during the summer of 2004. Journal of
 1150 Geophysical Research: Atmospheres, 110(D21).

1151 Makar, P.A., Gong, W., Milbrandt, J., Hogrefe, C., Zhang, Y., Curci, G., Žabkar, R., Im, U.,
 1152 Balzarini, A., Baró, R. and Bianconi, R., 2015a. Feedbacks between air pollution and weather.
 1153 Part 1: Effects on weather. Atmospheric Environment, (115), pp.442-469.

1154 Makar, P.A., Gong, W., Hogrefe, C., Zhang, Y., Curci, G., Žabkar, R., Milbrandt, J., Im, U.,
 1155 Balzarini, A., Baró, R. and Bianconi, R., 2015b. Feedbacks between air pollution and weather,
 1156 part 2: effects on chemistry. Atmospheric environment, 115, pp.499-526.

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Indent: Left: 0 cm, Automatically adjust right indent when grid is defined, Space After: 10 pt, Adjust space between Latin and Asian text, Adjust space between Asian text and numbers

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: (Asian) +Body Asian (SimSun), 小四

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

1157 Menon, S. H. J.; Nazarenko, N.; Luo, Y. (2002). "Climate Effects of Black Carbon Aerosols in
1158 China and India." Science.

1159 Nenes, A., Pandis, S.N. and Pilinis, C., 1998. ISORROPIA: A new thermodynamic equilibrium
1160 model for multiphase multicomponent inorganic aerosols. Aquatic geochemistry, 4(1), pp.123-
1161 152.

1162 Odum, J.R., Huffmann, T., Bowman, F., Collins, D., Flagan, R.C., Seinfeld, J.H.,
1163 1996.Gas/Particle partitioning and secondary organic aerosol yields. Environ. Sci. Technol. 30,
1164 2580-2585.

1165 Peters-Lidard, C. D., E. M. Kemp, T. Matsui, J.A. Santanello Jr., S.V. Kumar, J.P. Jacob, T.
1166 Clune, W.-K. Tao, M. Chin, A. Hou, J.L. Case, D. Kim, K.-M. Kim, W. Lau, Y. Liu, J. Shi, D.
1167 Starr, Q. Tan, Z. Tao, B.F. Zaitchik, B. Zavodsky, S.Q. Zhang, and M. Zupanski, Integrated
1168 modeling of aerosol, cloud, precipitation and land processes at satellite-resolved scales.
1169 Environmental Modeling & Software, 67, 149-159, doi:10.1016/j.envsoft.2015.01.007, 2015.

1170 Peters, M.D., Kreidenweis, S.M., 2007. A single parameter representation of hygroscopic
1171 growth and cloud condensation nucleus activity. Atmos. Chem. Phys. 7, 1961-1971.

1172 Ramanathan, V. C., G. (2008). "Global and regional climate changes due to black carbon."
1173 Nature Geoscience1(4): 221-227.

1174 San José, R., Pérez, J.L., Balzarini, A., Baró, R., Curci, G., Forkel, R., Galmarini, S., Grell, G.,
1175 Hirtl, M., Honzak, L. and Im, U., 2015. Sensitivity of feedback effects in CBMZ/MOSAIC
1176 chemical mechanism. Atmospheric Environment, 115, pp.646-656.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

1177 [Schuster, G. L., et al. \(2006\). "Angstrom exponent and bimodal aerosol size distributions."](#)
 1178 [Journal of Geophysical Research](#)111(D7).

1179 ▲

1180 [Seinfeld, J.H. and Pandis, S.N., 2016. Atmospheric chemistry and physics: from air pollution to](#)
 1181 [climate change. John Wiley & Sons.](#)

1182 [Stockwell, W. R., et al. \(1997\). "A new mechanism for regional atmospheric chemistry](#)
 1183 [modeling." Journal of Geophysical Research: Atmospheres](#)102(D22): 25847-25879.

1184 [Stoiber, R. E. W., S.N.; Huebert, B. \(1987\). "Annual contribution of sulfur dioxide to the](#)
 1185 [atmosphere by volcanoes." Journal of Volcanology and Geothermal Research](#)33: 1-8.

1186 [Streets, D. G., et al. \(2003\). "Biomass burning in Asia: Annual and seasonal estimates and](#)
 1187 [atmospheric emissions." Global Biogeochemical Cycles](#)17(4): n/a-n/a.

1188 [Tao, Z., J. A. Santanello, M. Chin, S. Zhou, Q. Tan, E. M. Kemp, and C. D. Peters-Lidard, Effect](#)
 1189 [of land cover on atmospheric processes and air quality over the continental United States – A](#)
 1190 [NASA Unified WRF \(NU-WRF\) model study. Atmospheric Chemistry & Physics, 13: 6207-](#)
 1191 [6226, doi: 10.5194/acp-13-6207-2013, 2013.](#)

1192 [Tao, Z., H. Yu, and M. Chin, Impact of transpacific aerosol on air quality over the United States:](#)
 1193 [A perspective from aerosol-cloud-radiation interactions. Atmospheric Environment, 125: 48-60,](#)
 1194 [doi:10.1016/j.atmosenv.2015.10.083, 2016.](#)

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Indent: First line: 0 cm, Automatically adjust right indent when grid is defined, Space After: 10 pt, Adjust space between Latin and Asian text, Adjust space between Asian text and numbers

Formatted: Indent: Left: 0 cm, Automatically adjust right indent when grid is defined, Space After: 10 pt, Adjust space between Latin and Asian text, Adjust space between Asian text and numbers

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1195 Tao, Z., H. Yu, and M. Chin. The role of aerosol-cloud-radiation interactions in regional air
1196 quality – A NU-WRF study over the United States. Atmosphere, 6, 1045-1068,
1197 doi:10.3390/atmos6081045, 2015.

1198 Tsimpidi et al., 2010. Evaluation of the volatility basis-set approach for the simulation of organic
1199 aerosol formation in the Mexico City metropolitan area. Atmos. Chem. Phys., 10, 525–546.
1200 Atmospheres102(D23): 28589-28612.

1201 Yu, S., Mathur, R., Pleim, J., Wong, D., Gilliam, R., Alapaty, K., Zhao, C. and Liu, X., 2013.
1202 Aerosol indirect effect on the grid-scale clouds in the two-way coupled WRF-CMAQ: model
1203 description, development, evaluation and regional analysis. Atmospheric Chemistry and Physics
1204 Discussion, p.25649,

1205 Wang et al., 2016. Persistent sulfate formation from London Fog to Chinese haze. PNAS,
1206 113(48), 13630–13635.

1207 Wang, J., et al. (2014). "Impact of aerosol–meteorology interactions on fine particle pollution
1208 during China’s severe haze episode in January 2013." Environmental Research Letters9(9):
1209 094002,

1210 Wang, T., et al. (2010). "Investigations on direct and indirect effect of nitrate on temperature and
1211 precipitation in China using a regional climate chemistry modeling system." Journal of
1212 Geophysical Research115.

1213 Wang, Z. Maeda, T.; Hayashi, M.; Hsiao, L.F.; Liu, K.Y. (2001). "A nested air quality
1214 prediction modeling system for urban and regional scales: application for high-ozone episode in
1215 Taiwan." Water, Air, & Soil Pollution130(1): 391-396

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, Font color: Auto, Pattern: Clear

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1216 [Xiao, H., et al. \(1997\). "Long-range transport of Sox and dust in East Asia during the PEM B](#)
1217 [Experiment." Journal of Geophysical Research;](#)

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1218 [Xin, J., et al. \(2015\). "The Campaign on Atmospheric Aerosol Research Network of China:](#)
1219 [CARE-China." Bulletin of the American Meteorological Society](#)96(7): 1137-1155.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1220 [Zaveri, R. A., et al. \(2008\). "Model for Simulating Aerosol Interactions and Chemistry](#)
1221 [\(MOSAIC\)." Journal of Geophysical Research](#)113(D13).

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1222 [Zaveri, R. A. and L. K. Peters \(1999\). "A new lumped structure photochemical mechanism for](#)
1223 [large-scale applications." Journal of Geophysical Research: Atmospheres](#)104(D23): 30387-
1224 [30415.](#)

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1225 [Zhang, Y., et al. \(2010\). "Simulating chemistry–aerosol–cloud–radiation–climate feedbacks over](#)
1226 [the continental U.S. using the online-coupled Weather Research Forecasting Model with](#)
1227 [chemistry \(WRF/Chem\)." Atmospheric Environment](#)44(29): 3568-3582.

Formatted: Font: 小四, No underline

Formatted: Font: 小四

1228 [Zheng et al., 2015. Heterogeneous chemistry: a mechanism missing in current models to explain](#)
1229 [secondary inorganic aerosol formation during the January 2013 haze episode in North China.](#)
1230 [Atmos. Chem. Phys., 15, 2031–204](#)

1231

1232 [Akimoto, H. \(2003\). "Global Air Quality and Pollution." Science, 302 \(5651\), 1716–1719.](#)

1233 [Archer-Nicholls, S., et al. \(2014\). "Gaseous chemistry and aerosol mechanism developments for](#)
1234 [version 3.5.1 of the online regional model, WRF-Chem." Geoscientific Model Development](#)7(6):
1235 [2557–2579.](#)

1236 [Baklanov, A., et al. \(2014\). "Online coupled regional meteorology chemistry models in Europe:](#)
1237 [current status and prospects." Atmospheric Chemistry and Physics](#)14(1): 317–398.

1240

1241 Baklanov, A., Brunner, D., Carmichael, G., Flemming, J., Freitas, S., Gauss, M., Hov, Ø.,
1242 Mathur, R., Schlünzen, K.H., Seigneur, C. and Vogel, B., 2017. Key issues for seamless
1243 integrated chemistry meteorology modeling. *Bulletin of the American Meteorological Society*,
1244 (2017).
1245 _____
1246 Bey, I., et al. (2001). "Global modeling of tropospheric chemistry with assimilated meteorology:
1247 Model description and evaluation." *Journal of Geophysical Research: Atmospheres*106(D19):
1248 23073-23095.
1249 _____
1250 Carmichael, G., et al. (2008). "MICS-Asia II: The model intercomparison study for Asia Phase II
1251 methodology and overview of findings." *Atmospheric Environment*42(15): 3468-3490.
1252 _____
1253 Carmichael, G. C., G.; Hayami, H.; Uno, I.; Cho, S.Y.; Engardt, M.; Kim, S.B.; Ichikawa, Y.;
1254 Ikeda, Y.; Woo, J.H.; Ueda, H.; Amann, M. (2002). "The MICS-Asia study: model
1255 intercomparison of long range transport and sulfur deposition in East Asia." *Atmospheric*
1256 *Environment*36: 175-199.
1257 _____
1258 Carmichael, G. P., L.R. (1984). "An Eulerian transport/transformation/removal model for SO₂
1259 and sulfate I. model development." *Atmospheric Environment*18(5): 937-951.
1260 _____
1261 Carmichael, G. P., L.R.; Kitada, T. (1986). "A second generation model for regional scale
1262 transport/chemistry/deposition." *Atmospheric Environment*20(1): 173-188.
1263 _____
1264 Carmichael, G. R., Peters, L.R.; Saylor, R. D. (1991). "The STEM-II regional scale acid
1265 deposition and photochemical oxidant model I. an overview of model development and
1266 applications." *Atmospheric Environment*25A(10): 2077-2090.
1267 _____
1268 Carmichael, G. R., et al. (2008). "Predicting air quality: Improvements through advanced
1269 methods to integrate models and measurements." *Journal of Computational Physics*227(7): 3540-
1270 3571.
1271 _____
1272 Carmichael, G. R., et al. (1998). "Tropospheric ozone production and transport in the springtime
1273 in east Asia." *Journal of Geophysical Research: Atmospheres*103(D9): 10649-10671.
1274 _____
1275 Carmichael, G. R. C., G.; Hayami, H.; Uno, I.; Cho, S.Y.; Engardt, M.; Kim, S.B.; Ichikawa, Y.;
1276 Ikeda, Y.; Woo, J.H.; Ueda, H.; Amann, M. (2002). "The MICS-Asia study: model
1277 intercomparison of long range transport and sulfur deposition in East Asia." *Atmospheric*
1278 *Environment*36: 175-199.
1279 _____
1280 Carter, W.P., 2000a. Documentation of the SAPRC-99 chemical mechanism for VOC reactivity
1281 assessment. Contract, 92(329), pp.95-308.
1282 _____
1283 Carter, W.P., 2000b. Implementation of the SAPRC-99 chemical mechanism into the models-3
1284 framework. Report to the United States Environmental Protection Agency, January, 29.
1285 _____

1286 Cheng, Y. F., Z., G.; Wei, C.; Mu, Q.; Zheng, B.; Wang, Z.; Gao, M.; Zhang, Q.; He, K.;
1287 Carmichael, G.; Poschl, U.; Su, Hang (2016). "Reactive nitrogen chemistry in aerosol water as a
1288 source of sulfate during haze events in China." Science Advances 2(e1601530).
1289 _____
1290 Chin, M. G., P.; Kinne, S.; Torres, O.; Holben, B.N.; Duncan, B.N.; Martin, R.V.; Logan, J.A.;
1291 Higurashi, A.; Nakajima, T. (2002). "Tropospheric aerosol optical thickness from the GOCART
1292 Model and Comparisons with satellite and sun photometer measurements." Journal of
1293 Atmospheric Sciences(59).
1294 _____
1295 Colarco, P., et al. (2010). "Online simulations of global aerosol distributions in the NASA
1296 GEOS 4 model and comparisons to satellite and ground-based aerosol optical depth." Journal of
1297 Geophysical Research115(D14).
1298 _____
1299 D'Almeida, G. A., P. Koepke, and E. P. Shettle (1991), Atmospheric Aerosols: Global
1300 Climatology and Radiative Characteristics, A. Deepak, Hampton, Va.
1301 _____
1302 Emery, C. T., E.; Yarwood, G. (2001). Enhanced meteorological modeling and performance
1303 evaluation for two Texas ozone episodes.
1304 _____
1305 Emmons, L. K. W., S.; Hess, P.G.; Lamarque, J.F.; Pfister G.G.; Fillmore, D.; Granier, C.;
1306 Guenther, A.; Kinnison, D.; Laeple, T.; Orlando, J.; Tie, X.; Tyndall, G.; Wiedinmyer, C.;
1307 Baughcum, S.L.; Kloster, S. (2010). "Description and evaluation of the Model for Ozone and
1308 Related chemical Tracers, version 4 (MOZART 4)." Geoscientific Model Development3: 43-67.
1309 _____
1310 Fu, J. S., N. C. Hsu, Y. Gao, K. Huang, C. Li, N. H. Lin, S. C. Tsay (2012). Evaluating the
1311 influences of biomass burning during 2006 BASE-ASIA: A regional chemical transport
1312 modeling. Atmospheric Chemistry and Physics, 12, 3837-3855.
1313 _____
1314 _____
1315 Gao, M., et al. (2016). "Response of winter fine particulate matter concentrations to emission and
1316 meteorology changes in North China." Atmospheric Chemistry and Physics16(18): 11837-
1317 11851.
1318 _____
1319 Gao, M., et al. (2016). "Improving simulations of sulfate aerosols during winter haze over
1320 Northern China: the impacts of heterogeneous oxidation by NO₂." Frontiers of Environmental
1321 Science & Engineering10(5).
1322 _____
1323 Gao, M., et al. (2016). "Modeling study of the 2010 regional haze event in the North China
1324 Plain." Atmospheric Chemistry and Physics16(3): 1673-1691.
1325 _____
1326 Gao, M., et al. (2015). "Health impacts and economic losses assessment of the 2013 severe haze
1327 event in Beijing area." Sci Total Environ511: 553-561.
1328 _____
1329 Gao, M., et al. (2017). "Estimates of Health Impacts and Radiative Forcing in Winter Haze in
1330 eastern China through constraints of surface PM_{2.5} predictions." Environ Sci Technol.
1331 _____

1332 Gery, M. W. W., G.Z.; Killus, J.P.; Dodge, M.C. (1989). "A photochemical kinetics mechanism
1333 for urban and regional scale computer modeling." Journal of Geophysical Research94(D10):
1334 12925-12956.
1335 _____
1336 Gong, D. Y. H., C.H. (2002). "<Gong and Ho, 2002.pdf>." Theoretical and Applied
1337 Climatology72: 1-9.
1338 _____
1339 Grell, G. A., et al. (2005). "Fully coupled "online" chemistry within the WRF model."
1340 Atmospheric Environment39(37): 6957-6975.
1341 _____
1342 Gutenther, A. K., T.; Harley, P.; Wiedinmyer, C.; Palmer, P.I.; Geron, C. (2006). "Estimates of
1343 global terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols
1344 from Nature)." Atmospheric Chemistry and Physics6: 3181-3210.
1345 _____
1346 Han, Zhiwei, Hiromasa Ueda, Kazuhide Matsuda, Renjian Zhang, Kimio Arai, Yutaka Kanai,
1347 Hisashi Hasome, 2004. Model study on particle size segregation and deposition during Asian
1348 dust events in March 2002, Journal of Geophysical Research, 109, D19205, doi:
1349 10.1029/2004jd004920.
1350 _____
1351 Han, Zhiwei. (2010). "Direct radiative effect of aerosols over East Asia with a Regional coupled
1352 Climate/Chemistry model." Meteorologische Zeitschrift,19(3): 287-298.
1353 _____
1354 Han Zhiwei, Jiawei Li, Xiangao Xia, Renjian Zhang, 2012. Investigation of direct radiative
1355 effects of aerosols in dust storm season over East Asia with an online coupled
1356 regional climate-chemistry-aerosol model. Atmospheric Environment, 54, 688-699.
1357 _____
1358 Han Zhiwei, Jiawei Li, Weidong Guo, Zhe Xiong, Wu Zhang, 2013. A study of dust radiative
1359 feedback on dust cycle and meteorology over East Asia by a coupled regional climate-chemistry-
1360 aerosol model. Atmospheric Environment, 68, 54-63.
1361 _____
1362 Han Zhiwei et al.,2016. Modeling organic aerosols over east China using a volatility basis set
1363 approach with aging mechanism in a regional air quality model. Atmospheric Environment 124,
1364 186-198.
1365 _____
1366 Hess, M., Koepke, P., Schuit, I., 1998. Optical properties of aerosols and clouds: the
1367 software package OPAC. Bull. Am. Meteorol. Soc. 79, 831-844.
1368 _____
1369 Holben, B. N., Eck, T.F., Slutsker, I., Tanre, D., Buis, J.P., Setzer, A., Vermote, E., Reagan, J.A.,
1370 Kaufman, Y.J., Nakajima, T. and Lavenu, F. (1998). "AERONET—A federated instrument
1371 network and data archive for aerosol characterization." Remote sensing of environment66(1): 1-
1372 16.
1373 _____
1374 Holloway, T., et al. (2008). "MICS Asia II: Impact of global emissions on regional air quality in
1375 Asia." Atmospheric Environment42(15): 3543-3561.
1376 _____

1377 [Huang, K. J. S., Fu, N. C., Hsu, Y., Gao, X., Dong, S. C., Tsay, Y. F., Lam \(2013\). Impact](#)
1378 [assessment of biomass burning on air quality in Southeast and East Asia during BASE-ASIA.](#)
1379 [Atmospheric Environment, 78, 291-302.](#)

1380

1381 [Jacobson, M. Z., 2001: Global direct radiative forcing due to multicomponent anthropogenic and](#)
1382 [natural aerosols. J. Geophys. Res., 106, 1551–1568.](#)

1383

1384 [Janssens-Maenhout, G., et al. \(2015\). "HTAP_v2.2: a mosaic of regional and global emission](#)
1385 [grid maps for 2008 and 2010 to study hemispheric transport of air pollution." Atmospheric](#)
1386 [Chemistry and Physics](#)15(19): 11411-11432.

1387

1388 [Kiehl, J.T., Briegleb, B.P., 1993. The relative roles of sulfate aerosols and greenhouse gases in](#)
1389 [climate forcing. Science 260, 311-314.](#)

1390

1391 [Kim, S. W., et al. \(2009\). "NO₂ columns in the western United States observed from space and](#)
1392 [simulated by a regional chemistry model and their implications for NO_x emissions." Journal of](#)
1393 [Geophysical Research](#)114(D11).

1394

1395 [Kim, Y. J. S., S.N.; Carmichael, G.R.; Riemer, N.; Stanier, C.O. \(2014\). "Modeled aerosol](#)
1396 [nitrate formation pathways during wintertime in the Great Lakes region of North America."](#)
1397 [Journal of Geophysical Research: Atmospheres](#)119: 12420-12445.

1398

1399 [Knote, C., et al. \(2015\). "Influence of the choice of gas-phase mechanism on predictions of key](#)
1400 [gaseous pollutants during the AQMEII phase-2 intercomparison." Atmospheric Environment](#)115:
1401 [553-568.](#)

1402

1403 [Lack, D. A., et al. \(2004\). "Seasonal variability of secondary organic aerosol: A global modeling](#)
1404 [study." Journal of Geophysical Research: Atmospheres](#)109(D3): n/a-n/a.

1405

1406 [Lelieveld, J., et al. \(2015\). "The contribution of outdoor air pollution sources to premature](#)
1407 [mortality on a global scale." Nature](#)525(7569): 367-371.

1408

1409 [Li, M., et al. \(2017\). "MIX: a mosaic Asian anthropogenic emission inventory under the](#)
1410 [international collaboration framework of the MICS-Asia and HTAP." Atmospheric Chemistry](#)
1411 [and Physics](#)17(2): 935-963.

1412

1413 [Menon, S. H., J.; Nazarenko, N.; Luo, Y. \(2002\). "Climate Effects of Black Carbon Aerosols in](#)
1414 [China and India." Science.](#)

1415

1416 [Odum, J.R., Huffmann, T., Bowman, F., Collins, D., Flagan, R.C., Seinfeld, J.H., 1996.](#)
1417 [Gas/Particle partitioning and secondary organic aerosol yields. Environ. Sci. Technol. 30, 2580-](#)
1418 [2585.](#)

1419

1420 [Peters-Lidard, C. D., E. M. Kemp, T. Matsui, J.A. Santanello Jr., S.V. Kumar, J.P. Jacob, T.](#)
1421 [Clune, W. K. Tao, M. Chin, A. Hou, J.L. Case, D. Kim, K. M. Kim, W. Lau, Y. Liu, J. Shi, D.](#)
1422 [Starr, Q. Tan, Z. Tao, B.F. Zaitchik, B. Zavadsky, S.Q. Zhang, and M. Zupanski, Integrated](#)

1423 modeling of aerosol, cloud, precipitation and land processes at satellite-resolved scales.
1424 *Environmental Modeling & Software*, 67, 149–159, doi:10.1016/j.envsoft.2015.01.007, 2015.
1425
1426 Petters, M.D., Kreidenweis, S.M., 2007. A single parameter representation of hygroscopic
1427 growth and cloud condensation nucleus activity. *Atmos. Chem. Phys.* 7, 1961–1971.
1428
1429 Ramanathan, V. C., G. (2008). "Global and regional climate changes due to black carbon."
1430 *Nature Geoscience*1(4): 221–227.
1431
1432 Schuster, G. L., et al. (2006). "Angstrom exponent and bimodal aerosol size distributions."
1433 *Journal of Geophysical Research*111(D7):
1434
1435 Seinfeld, J. H., and S. N. Pandis. (2006). *Atmospheric chemistry and physics*.
1436
1437 Shrivastava, M., et al. (2011). "Modeling organic aerosols in a megacity: comparison of simple
1438 and complex representations of the volatility basis set approach." *Atmospheric Chemistry and*
1439 *Physics*11(13): 6639–6662.
1440
1441 Stockwell, W. R., et al. (1997). "A new mechanism for regional atmospheric chemistry
1442 modeling." *Journal of Geophysical Research: Atmospheres*102(D22): 25847–25879.
1443
1444 Stoiber, R. E. W., S.N.; Huebert, B. (1987). "Annual contribution of sulfur dioxide to the
1445 atmosphere by volcanoes." *Journal of Volcanology and Geothermal Research*33: 1–8.
1446
1447 Streets, D. G., et al. (2003). "Biomass burning in Asia: Annual and seasonal estimates and
1448 atmospheric emissions." *Global Biogeochemical Cycles*17(4): n/a–n/a.
1449
1450 Sudo, K., et al. (2002). "CHASER: A global chemical model of the troposphere I. Model
1451 description." *Journal of Geophysical Research: Atmospheres*107(D17): ACH 7–1–ACH 7–20.
1452
1453 Tao, Z., H. Yu, and M. Chin, Impact of transpacific aerosol on air quality over the United States:
1454 A perspective from aerosol–cloud–radiation interactions. *Atmospheric Environment*, 125: 48–60,
1455 doi:10.1016/j.atmosenv.2015.10.083, 2016.
1456
1457 Tao, Z., H. Yu, and M. Chin, The role of aerosol–cloud–radiation interactions in regional air
1458 quality—A NU-WRF study over the United States. *Atmosphere*, 6, 1045–1068,
1459 doi:10.3390/atmos6081045, 2015.
1460
1461 Tao, Z., J. A. Santanello, M. Chin, S. Zhou, Q. Tan, E. M. Kemp, and C. D. Peters-Lidard, Effect
1462 of land cover on atmospheric processes and air quality over the continental United States—A
1463 NASA Unified-WRF (NU-WRF) model study. *Atmospheric Chemistry & Physics*, 13: 6207–
1464 6226, doi: 10.5194/acp-13-6207-2013, 2013.
1465
1466 Tsimpidi et al., 2010. Evaluation of the volatility basis set approach for the simulation of organic
1467 aerosol formation in the Mexico City metropolitan area. *Atmos. Chem. Phys.*, 10, 525–546.
1468

1469 Wang et al., 2016. Persistent sulfate formation from London Fog to Chinese haze. *PNAS*,
1470 113(48), 13630–13635.
1471
1472 Wang, J., et al. (2014). "Impact of aerosol–meteorology interactions on fine particle pollution
1473 during China's severe haze episode in January 2013." *Environmental Research Letters*9(9):
1474 094002.
1475 _____
1476 Wang, T., et al. (2010). "Investigations on direct and indirect effect of nitrate on temperature and
1477 precipitation in China using a regional climate chemistry modeling system." *Journal of*
1478 *Geophysical Research*115.
1479 _____
1480 Wang, Z., Maeda, T.; Hayashi, M.; Hsiao, L.F.; Liu, K.Y. (2001). "A nested air quality
1481 prediction modeling system for urban and regional scales: application for high-ozone episode in
1482 Taiwan." *Water, Air, & Soil Pollution*130(1): 391–396.
1483 _____
1484 Xiao, H., et al. (1997). "Long-range transport of SO_x and dust in East Asia during the PEM-B
1485 Experiment." *Journal of Geophysical Research: Atmospheres*102(D23): 28589–28612.
1486 _____
1487 Xin, J., et al. (2015). "The Campaign on Atmospheric Aerosol Research Network of China:
1488 CARE-China." *Bulletin of the American Meteorological Society*96(7): 1137–1155.
1489 _____
1490 Zaveri, R. A., et al. (2008). "Model for Simulating Aerosol Interactions and Chemistry
1491 (MOSAIC)." *Journal of Geophysical Research*113(D13).
1492 _____
1493 Zaveri, R. A. and L. K. Peters (1999). "A new lumped structure photochemical mechanism for
1494 large-scale applications." *Journal of Geophysical Research: Atmospheres*104(D23): 30387–
1495 30415.
1496 _____
1497 Zhang, Y., et al. (2010). "Simulating chemistry–aerosol–cloud–radiation–climate feedbacks over
1498 the continental U.S. using the online coupled Weather Research Forecasting Model with
1499 chemistry (WRF/Chem)." *Atmospheric Environment*44(29): 3568–3582.
1500 _____
1501 Zhang, Y. C., Y.; Sarwar, G.; Schere, K. (2012). "Impact of gas-phase mechanisms on Weather
1502 Research Forecasting Model with Chemistry (WRF/Chem) predictions: Mechanism
1503 implementation and comparative evaluation." *Journal of Geophysical Research*117(D01301).
1504 _____
1505 Zheng et al., 2015. Heterogeneous chemistry: a mechanism missing in current models to explain
1506 secondary inorganic aerosol formation during the January 2013 haze episode in North China.
1507 *Atmos. Chem. Phys.*, 15, 2031–204
1508
1509
1510
1511
1512
1513
1514

1515
 1516
 1517
 1518
 1519
 1520
 1521
 1522
 1523
 1524
 1525
 1526
 1527
 1528
 1529
 1530
 1531
 1532
 1533
 1534
 1535
 1536
 1537
 1538
 1539
 1540
 1541
 1542
 1543
 1544
 1545
 1546
 1547
 1548
 1549
 1550

<u>Models</u>	<u>M1: WRF-Chem</u>	<u>M2: WRF-Chem</u>	<u>M3: NU-WRF</u>	<u>M4: NU-WRF</u>	<u>M5: RIEMS-Chem</u>	<u>M6: RegCCMS</u>	<u>M7: WRF-CMAQ</u>
<u>Modelling Group</u>	<u>Pusan National University</u>	<u>University of Iowa</u>	<u>USRA/NASA</u>	<u>USRA/NASA</u>	<u>Institute of Atmospheric Physics</u>	<u>Nanjing University</u>	<u>University of Tennessee</u>
<u>Grid Resolution</u>	<u>45km</u>	<u>50km</u>	<u>45km</u>	<u>15km</u>	<u>60km</u>	<u>50km</u>	<u>45km</u>

<u>Vertical Layers</u>	<u>40 layers to 50mb</u>	<u>27 layers to 50mb</u>	<u>60 layers to 20mb</u>	<u>60 layers to 20mb</u>	<u>16 layers to 100mb</u>	<u>18 layers to 50mb</u>	
<u>Gas phase chemistry</u>	<u>RACM-ESRL</u>	<u>CBMZ</u>	<u>RADM2</u>	<u>RADM2</u>	<u>CBM4</u>	<u>CBM4</u>	<u>SAPRC99</u>
<u>Aerosols</u>	<u>MADE/SOR GAM modal scheme</u>	<u>MOSAIC-8bin</u>	<u>GOCART bulk scheme</u>	<u>GOCART bulk scheme</u>	<u>Sulfate, nitrate, ammonium, BC, OC, SOA, 5 bins of soil dust, and 5 bins of sea salt modal scheme</u>	<u>Sulfate, nitrate, ammonium, BC and POC bulk scheme</u>	<u>AE06 modal scheme</u>
<u>Chemical Boundary Conditions</u>	<u>Climatological data from NALROM</u>	<u>MOZART</u>	<u>MOZART GOCART</u>	<u>MOZART GOCART</u>	<u>GEOS-Chem</u>	<u>Climatological data</u>	<u>GEOS-Chem</u>
<u>Meteorological Boundary Conditions</u>	<u>NCEP FNL</u>	<u>NCEP FNL</u>	<u>NASA MERRA</u>	<u>NASA MERRA</u>	<u>NCEP FNL</u>	<u>NCEP-NCAR</u>	<u>NCEP FNL</u>
<u>BVOC emissions</u>	<u>prescribed</u>	<u>Internal calculation</u>	<u>Internal calculation</u>	<u>Internal calculation</u>	<u>prescribed</u>	<u>No BVOC</u>	<u>Internal calculation</u>
<u>Dust</u>	<u>NA</u>	<u>EROD data</u>	<u>GOCART dust</u>	<u>GOCART dust</u>	<u>Han et al. (2004)</u>	<u>NA</u>	<u>No</u>
<u>Microphysics</u>	<u>Lin scheme</u>	<u>Morrison double-moment</u>	<u>GCE (Goddard Cumulus Ensemble)</u>	<u>GCE</u>	<u>Reisner mixed phase</u>		
<u>Longwave radiation</u>	<u>RRTMG</u>	<u>RRTMG</u>	<u>Goddard</u>	<u>Goddard</u>	<u>CCM3</u>		<u>RRTM</u>
<u>Shortwave radiation</u>	<u>RRTMG</u>	<u>RRTMG</u>	<u>Goddard</u>	<u>Goddard</u>	<u>Revised CCM3</u>		<u>Goddard</u>
<u>Boundary Layer Cu physics</u>	<u>Yonsei University Grell 3D</u>	<u>Yonsei University Grell 3D</u>	<u>YSU Grell 3D</u>	<u>YSU Grell 3D</u>	<u>MRF Grell 3D</u>	<u>*BATS</u>	
<u>Surface physics</u>	<u>Thermal diffusion</u>	<u>Unified Noah</u>	<u>Unified Noah</u>	<u>Unified Noah</u>	<u>BATS</u>		
<u>Aerosol-radiation</u>	<u>Yes</u>	<u>Yes</u>	<u>Yes</u>	<u>Yes</u>	<u>Yes</u>	<u>Yes</u>	<u>Yes</u>
<u>Aerosol-microphysics</u>	<u>Yes</u>	<u>Yes</u>	<u>Yes</u>	<u>Yes</u>	<u>Yes</u>	<u>Yes</u>	<u>No</u>
<u>Mixing state</u>	<u>Internal mixing</u>	<u>Internal mixing</u>	<u>Internal mixing</u>	<u>Internal mixing</u>	<u>Internal mixing among inorganic aerosols and BC and OC, and external mixing between dust, sea-salt and other aerosols</u>	<u>External mixing</u>	<u>Internal mixing</u>

1551 Table 1 Participating models in Topic 3

Models	M1: WRF-Chem1	M2: WRF-Chem2	M3: NU-WRF1	M4: NU-WRF2	M5: RIEMS-Chem	M6: RegCCMS	M7: WRF-CMAQ
Modelling Group	Pusan National University	University of Iowa	USRA/NASA A	USRA/NASA	Institute of Atmospheric Physics	Nanjing University	University of Tennessee
Grid Resolution	45km	50km	45km	15km	60km	50km	45km
Vertical Layers	40 layers to 50mb	27 layers to 50mb	60 layers to 20mb	60 layers to 20mb	16 layers to 100mb	18 layers to 50mb	
Gas phase chemistry	RACM	CBMZ	RADM2	RADM2	CBM4	CBM4	SAPRC99
Aerosols	MADE	MOSAIC-8bin	GOCART	GOCART	Sulfate, nitrate, ammonium, BC, OC, SOA, 5 bins of soil dust, and 5 bins of sea salt	Sulfate, nitrate, ammonium, BC and POC	AE06
Chemical Boundary Conditions	Climatological data from NALROM	MOZART	MOZART GOCART	MOZART GOCART	GEOS-Chem	Climatological data	GEOS-Chem

Participating models in Topic 3

M1: WRF-Chem v3.7.1; M2: WRF-Chem v3.5.1; M7: WRFv3.4.1&CMAQv5.0.2 NU-WRF v7lis7-3.5.1-p3

Table 2 CARE-Chine network sites
67

ID	Site name	Characteristics	Longitude	Latitude
1	Beijing	AOD	116.37	39.97
2	Tianjin	Air quality*	117.21	39.08
3	Shijiazhuang	Air quality	114.53	38.03
4	Xianghe	Air quality	116.96	39.75
5	Xinglong	Air quality	117.58	40.39
6	Beijing Forest	AOD	115.43	39.97
7	Baoding	AOD	115.51	38.87
8	Cangzhou	AOD	116.80	38.28
9	Shenyang	AOD	123.63	41.52
10	Jiaozhou Bay	AOD	120.18	35.90

*Air quality: surface PM_{2.5}, PM₁₀, SO₂, NO_x, CO, O₃

Table 3 Performance Statistics of Meteorology Variables (RMSE and MBE units: degree for T2; g/kg for Q2; m/s for WS10; W/m² for SWDOWN)

Metrics	Models	T2	Q2	WS10	SWDOWN	SWDOWN
					South	North
RMSE	M1	0.64	0.14	2.04	86.32	69.39
	M2	0.68	0.10	0.95	96.71	72.76
	M3	2.34	0.16	1.16	60.34	59.56
	M4	2.90	0.43	1.44	100.34	74.89
	M5	2.97	0.46	0.91	91.06	65.27
	M6	3.57	0.76	2.48	85.63	222.00
	M7	2.05	0.17	0.22	158.10	218.67
	Ensemble	1.81	0.10	1.28	81.96	62.51
MBE	M1	-0.19	0.02	2.01	66.58	59.94
	M2	-0.60	-0.01	0.91	83.88	62.38
	M3	-2.18	-0.04	1.11	36.44	47.74
	M4	-2.09	0.11	1.40	26.78	33.59
	M5	-2.73	0.43	0.74	49.06	51.00
	M6	-3.06	-0.56	2.37	-0.49	-202.26
	M7	-2.02	-0.12	0.15	145.24	159.02
	Ensemble	-1.71	-0.02	1.25	65.54	36.37
NMB (%)	M1	-0.07%	0.19%	17.58%	14.61%	13.34%
	M2	-0.21%	-0.12%	7.94%	18.41%	13.88%
	M3	-0.79%	-0.34%	9.73%	8.00%	10.63%
	M4	-0.75%	0.95%	12.26%	5.88%	7.48%
	M5	-0.98%	3.65%	6.45%	10.77%	11.35%
	M6	-1.10%	-4.77%	20.73%	-0.11%	-45.02%
	M7	-0.72%	-1.05%	1.31%	31.88%	35.39%
	Ensemble	-0.61%	-0.14%	10.98%	14.38%	8.10%

Table 4 Performance Statistics of Air Pollutants at the CARE-China sites (RMSE and MBE units: ppbv for gases and $\mu\text{g}/\text{m}^3$ for PM)

Metrics	Models	SO ₂	NO _x	O ₃	PM _{2.5}	PM ₁₀		SO ₂	NO _x	O ₃	PM _{2.5}	PM ₁₀
r	M1	0.76	0.60	0.46	0.85	0.76	MBE	-17.14	-5.53	-1.54	55.69	30.70
	M2	0.77	0.65	0.48	0.90	0.85		2.10	33.41	2.53	48.44	12.94
	M3	0.69	0.66	0.39	0.85	0.68		-15.89	-8.00	23.93	8.13	-19.92
	M4	0.67	0.61	0.42	0.88	0.73		-9.98	0.28	24.49	23.12	-3.23
	M5	0.72	0.73	0.39	0.91	0.84		-9.69	64.29	-5.30	1.68	-52.49
	M6	0.62	0.48	-	-	-		-27.53	-29.98	-	-	-
	M7	0.57	0.58	0.48	0.82	0.77		-25.56	7.85	-3.09	43.59	-21.00
Ensemble	0.79	0.71	0.51	0.94	0.87	-14.81	8.90	6.84	30.11	-8.83		
RMSE	M1	27.63	33.51	6.40	73.37	79.06	NMB (%)	-14.05	-5.41	7.37	63.57	18.93
	M2	21.00	66.30	8.15	72.44	80.72		12.13	69.58	39.87	54.07	6.38
	M3	29.50	36.87	24.76	47.20	78.21		-10.44	-6.26	306.33	9.67	-12.41
	M4	26.86	36.10	25.34	49.13	72.25		0.31	4.51	316.99	27.03	-1.78
	M5	32.17	87.48	7.90	45.32	81.00		6.83	127.45	-38.49	0.52	-32.94
	M6	33.95	48.62	-	-	-		-51.28	-48.59	-	-	-
	M7	34.75	35.88	6.89	64.25	70.19		-37.87	18.32	-7.78	48.92	-12.78
Ensemble	24.10	29.12	8.86	45.25	56.65	-13.48	22.80	104.04	33.96	-5.77		
MFB (%)	M1	-17.32	5.26	-5.06	64.34	21.98	MFE (%)	53.73	43.79	54.54	69.92	41.95
	M2	9.09	32.82	19.88	51.18	3.44		43.18	73.39	60.79	59.87	39.35
	M3	-12.96	4.52	113.60	32.67	-4.62		57.87	46.69	113.60	50.10	36.83
	M4	1.53	15.34	114.35	45.27	6.07		46.30	48.13	114.35	55.03	34.72
	M5	-20.24	67.25	-62.65	16.88	-35.15		63.69	72.07	80.92	48.17	45.09
	M6	-77.13	-56.89	-	-	-		84.21	69.66	-	-	-
	M7	-46.67	21.80	-19.50	57.19	-7.02		72.35	49.18	60.64	66.27	35.83
Ensemble	-14.17	26.41	62.86	50.61	3.12	43.13	42.94	71.14	55.86	28.05		

Table 5 Performance Statistics of Air Pollutants at the EANET sites (RMSE and MBE units: ppbv for gases and $\mu\text{g}/\text{m}^3$ for PM)

Metrics	Models	SO₂	NO_x	O₃	PM₁₀		SO₂	NO_x	O₃	PM₁₀
r	M1	0.57	0.64	0.14	0.59		-0.68	0.68	-6.16	-21.03
	M2	0.59	0.45	0.30	0.75		-0.45	-0.39	5.50	3.12
	M3	0.50	0.55	0.26	0.51		-0.37	-0.21	3.67	3.55
	M4	0.45	0.55	0.25	0.49		-0.57	-0.61	4.28	2.96
	M5	0.58	0.54	0.01	0.03		-0.57	1.28	4.67	3.77
	M6	0.33	0.24	-	-	MBE	0.32	-1.68	-	-
	M7	0.53	0.49	0.38	0.55		-0.03	0.64	-1.89	-15.75
	Ensemble	0.60	0.66	0.32	0.59		-0.34	-0.07	1.68	-3.89
NMB (%)	M1	-46.45	41.49	-15.03	-82.29		1.18	1.37	8.23	23.39
	M2	-29.64	-29.75	13.47	18.90		1.01	1.35	7.29	10.01
	M3	-25.42	-17.75	9.01	19.46		1.02	1.02	6.44	13.71
	M4	-39.63	-35.84	10.47	16.95	RMSE	1.14	0.97	6.35	13.78
	M5	-34.23	38.50	11.38	31.80		1.27	2.75	12.27	23.10
	M6	12.63	-93.57	-	-		1.38	1.85	-	-
	M7	17.42	31.47	-4.71	-56.18		1.04	1.57	6.52	18.76

Ensemble	-20.76	-10.79	4.10	-8.56	0.96	0.79	4.98	11.69
-----------------	--------	--------	------	-------	------	------	------	-------

Table 6 Performance Statistics of AOD

Metrics	Models	M1	M2	M3	M4	M5	M6	M7	Ensemble
R	North	0.63	0.74	0.57	0.51	0.68	0.36	0.71	0.77
	China								
	All	0.60	0.65	0.46	0.42	0.53	0.33	0.64	0.75
MBE	North	-0.25	-0.10	-0.09	-0.07	-0.13	-0.21	-0.05	-0.03
	China								
	All	-0.18	-0.02	-0.01	-0.01	-0.01	-0.11	0.00	-0.12
NMB (%)	North	-71.25	-23.28	-12.63	-9.59	-28.34	-59.19	-2.70	-30.17
	China								
	All	-74.94	-30.69	-25.68	-23.64	-28.24	-55.38	-21.12	-28.91
RMSE	North	0.35	0.20	0.26	0.28	0.24	0.36	0.22	0.22
	China								
	All	1.16	1.13	1.15	1.15	1.15	1.17	1.14	0.20

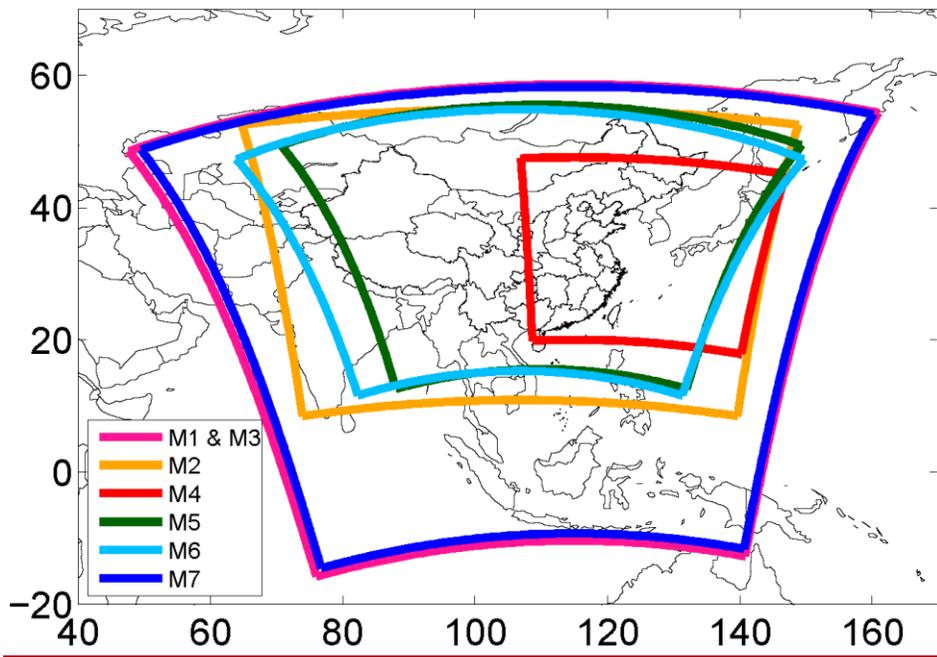
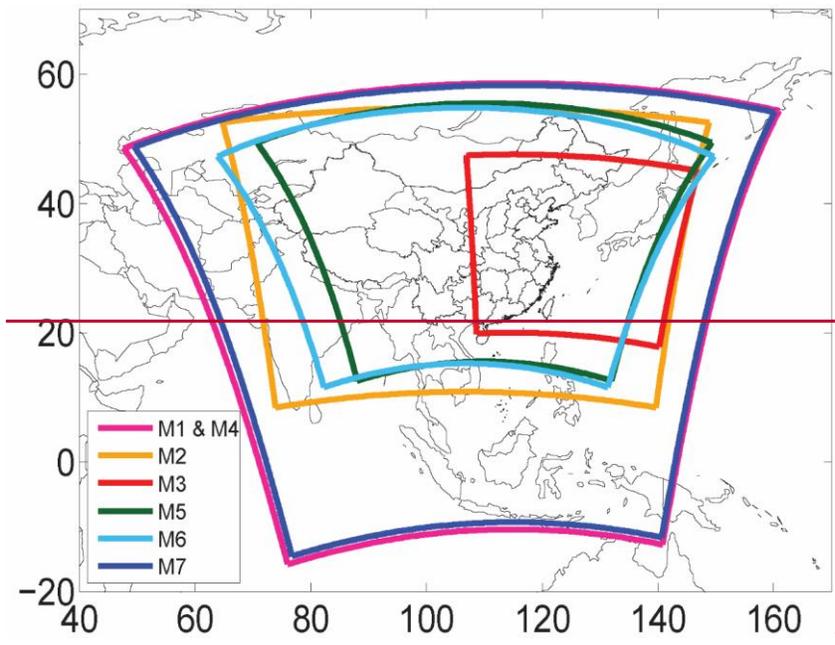


Figure 1. MICS-ASIA III Topic 3 modeling domains (descriptions of each model are documented in Table 1) M1: WRF-Chem 45km; M2: WRF-Chem 50km; M3: NU WRF 45km; M4: NU-WRF 15km; M5: RIEMS-IAP 60km; RegCCMS 50km; WRF-CMAQ 45km

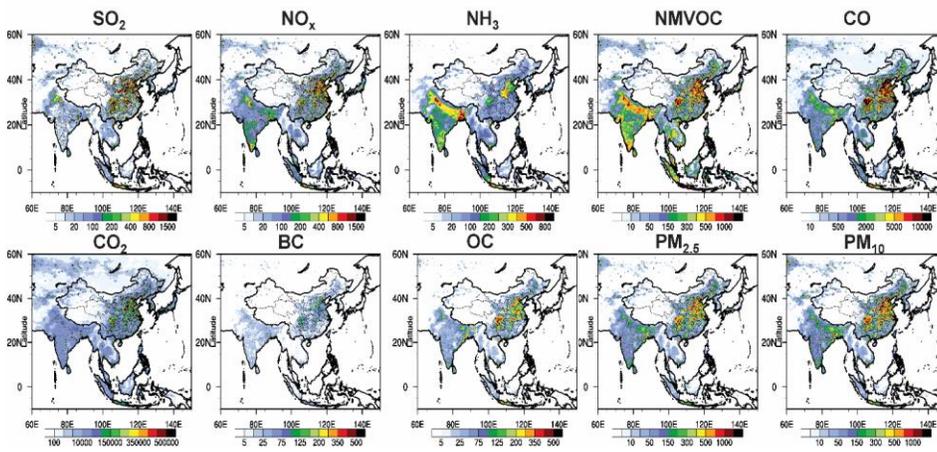


Figure 2. MIX emission inventory for January 2010 (Mg/month/grid)

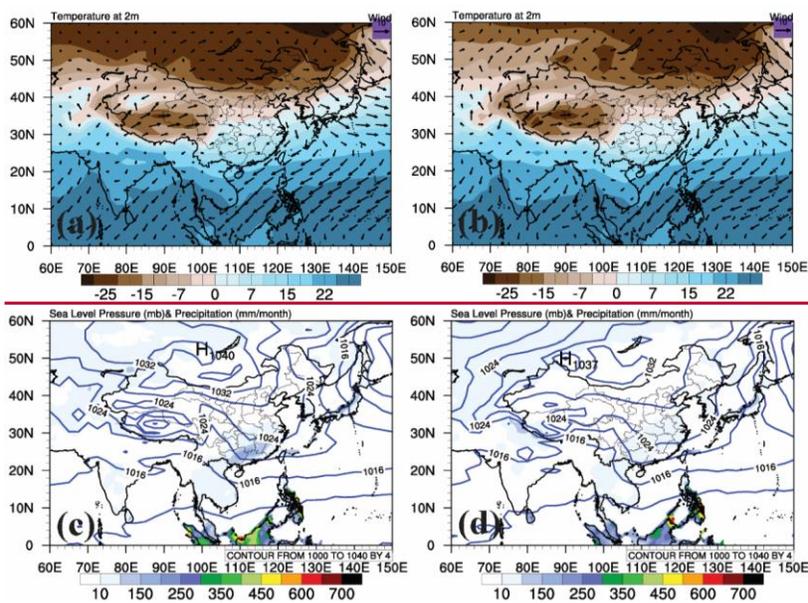


Figure 3. Monthly mean temperature at 2m, winds at 10m, total precipitation and sea level pressure for January 2010 (a,c) and January 2013 (b,d)

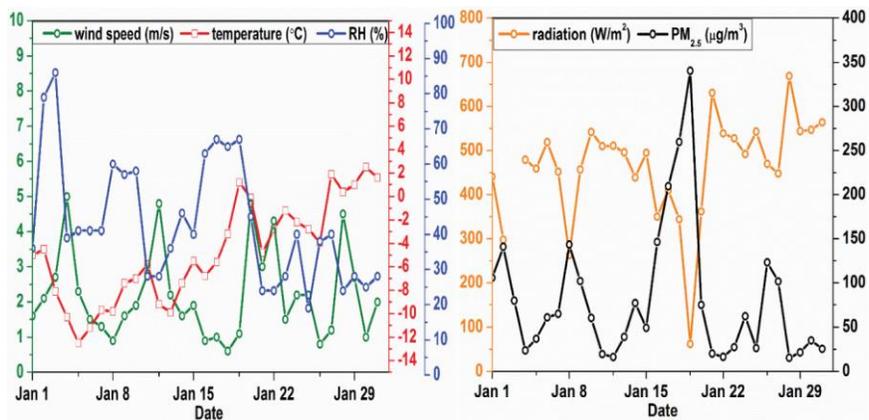


Figure 4. Observed near surface daily meteorological variables and PM_{2.5} concentrations in Beijing for January 2010

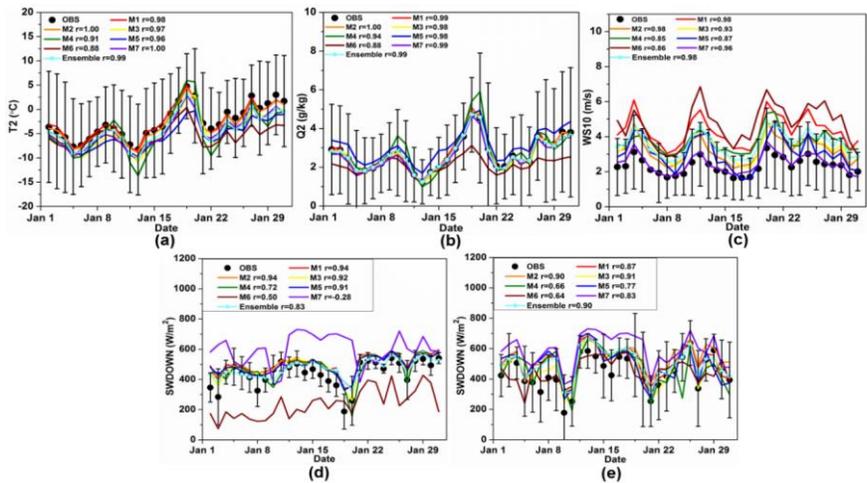


Figure 54. Comparisons between simulated and observed near surface temperature (a), water vapor mixing ratio (b), and wind speeds (c) (T2, Q2, and WS10), downward shortwave radiation in North China (d) and South China (e) (spatial daily values are averaged over measurements shown in S4 and S5; the error bars show the standard deviation of values over the measurement sites)

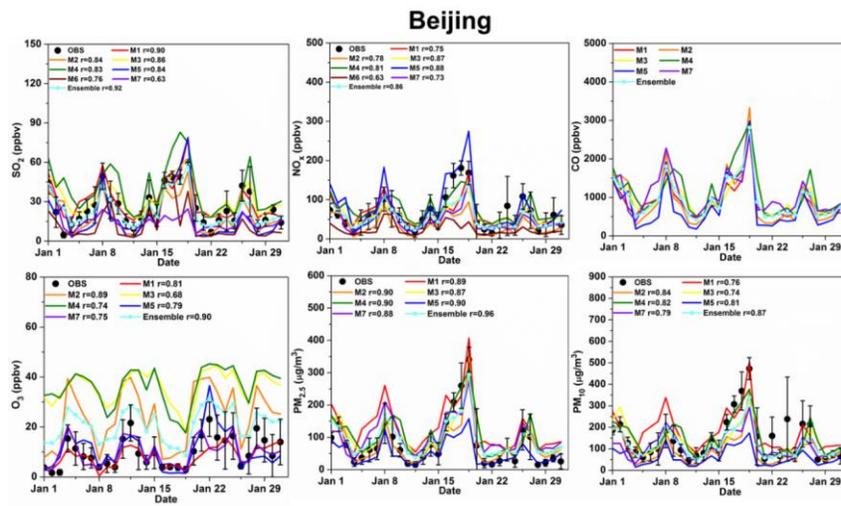


Figure 65. Comparisons between simulated and observed daily air pollutants (SO_2 , NO_x , CO , O_3 , $\text{PM}_{2.5}$ and PM_{10}) at the Beijing CARE-China site

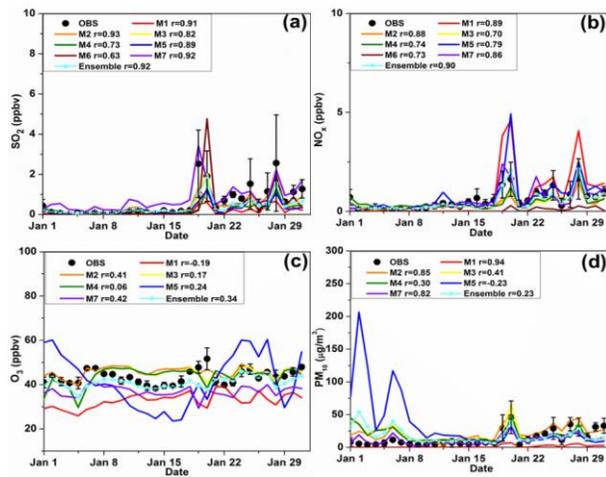


Figure 76. Comparisons between simulated and observed daily air pollutants (SO_2 , NO_x , O_3 , and PM_{10}) at the Rishiri EANET sites

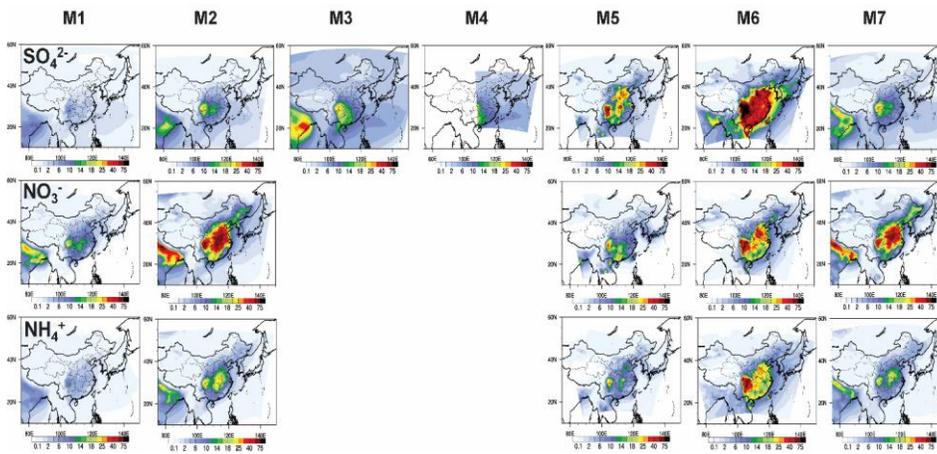


Figure 87. Simulated monthly concentrations of major PM_{2.5} components (μg/m³) for January 2010 from all participating models

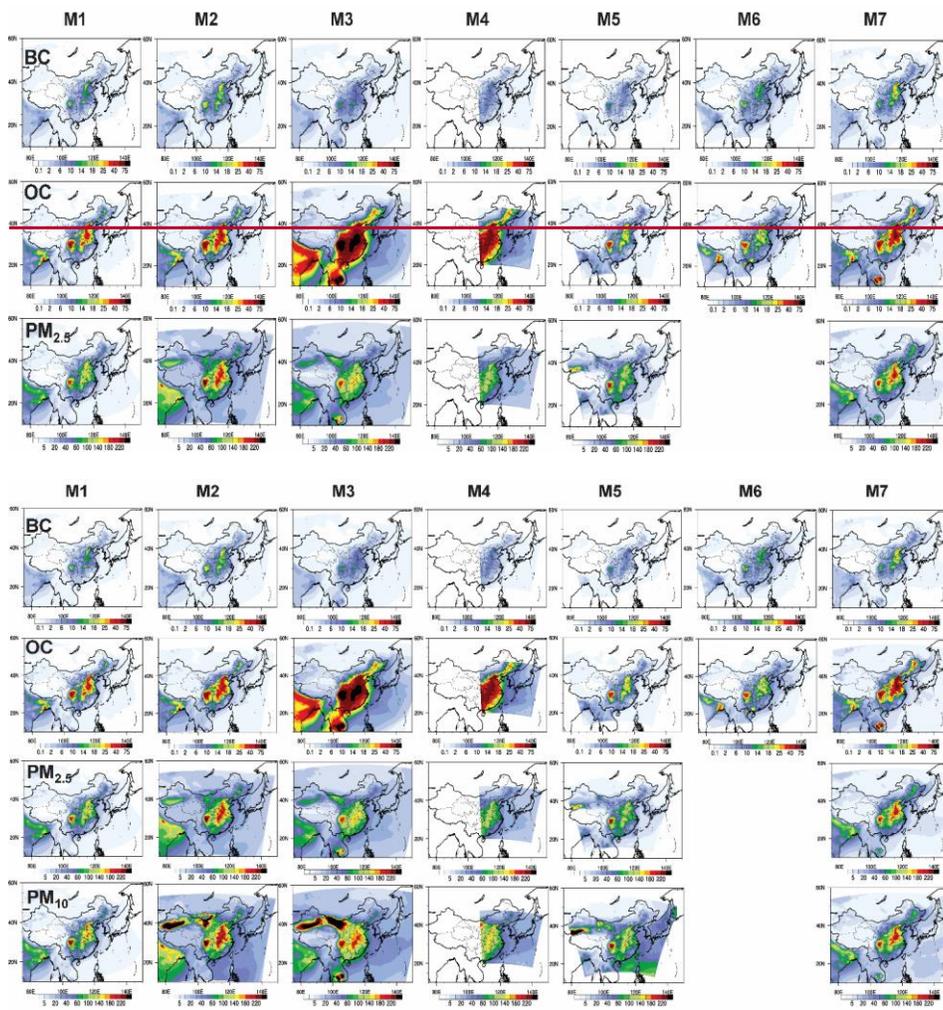


Figure 98. Simulated monthly concentrations of PM_{2.5} and major PM_{2.5} components ($\mu\text{g}/\text{m}^3$) for January 2010 from all participating models

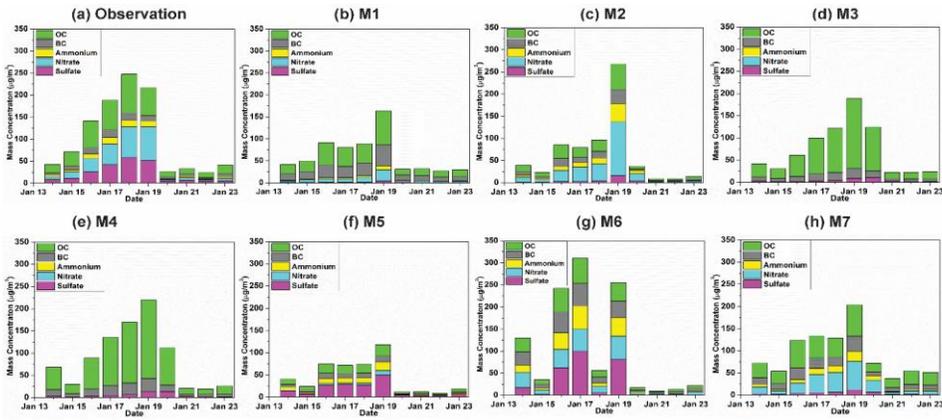


Figure 109. Observed and simulated daily mean concentrations of major PM_{2.5} chemical components in the urban Beijing site

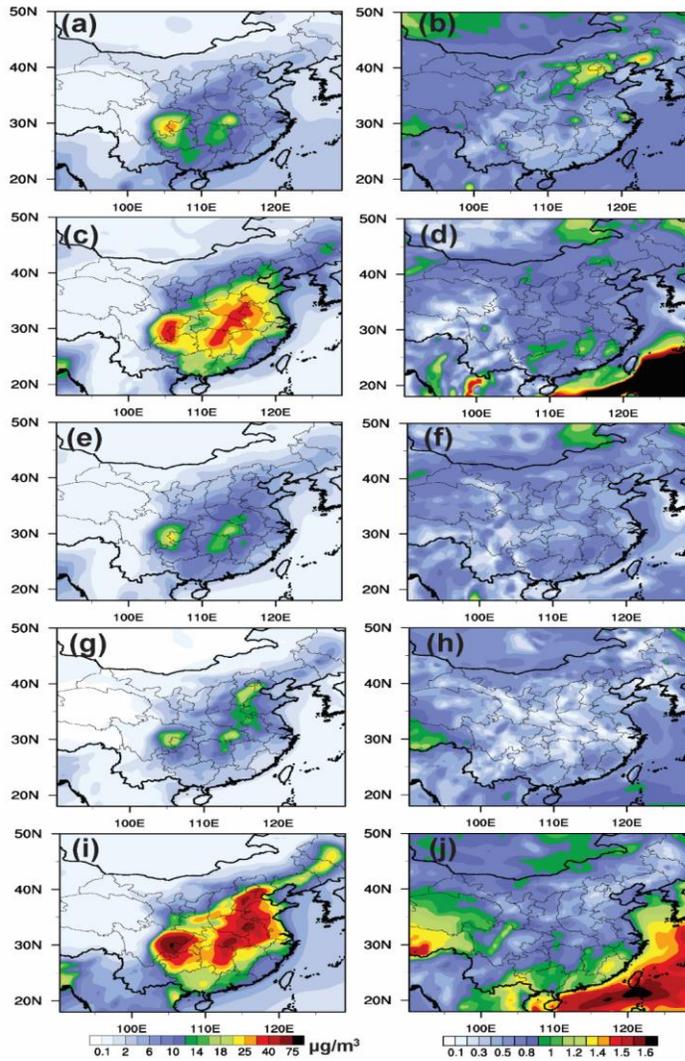


Figure 4-10. The ensemble mean monthly averaged near-surface distributions of $PM_{2.5}$ compositions for January 2010 (sulfate (a), nitrate (c), ammonium (e), BC (g), and OC (i)), along with the spatial distribution of the coefficient of variation ((b), (d), (f), (h), and (j), standard deviation divided by the average)

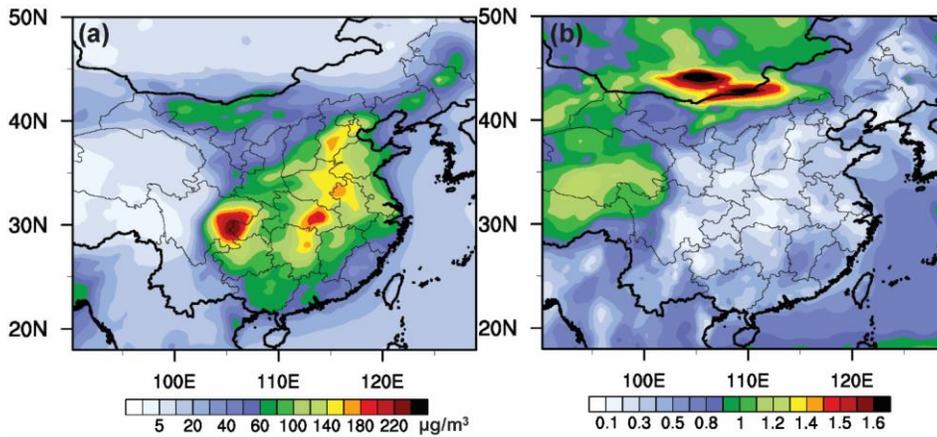


Figure 12.11. The ensemble mean monthly averaged near-surface distributions of $PM_{2.5}$ for January 2010 (a), along with the spatial distribution of the coefficient of variation (b, standard deviation divided by the average)

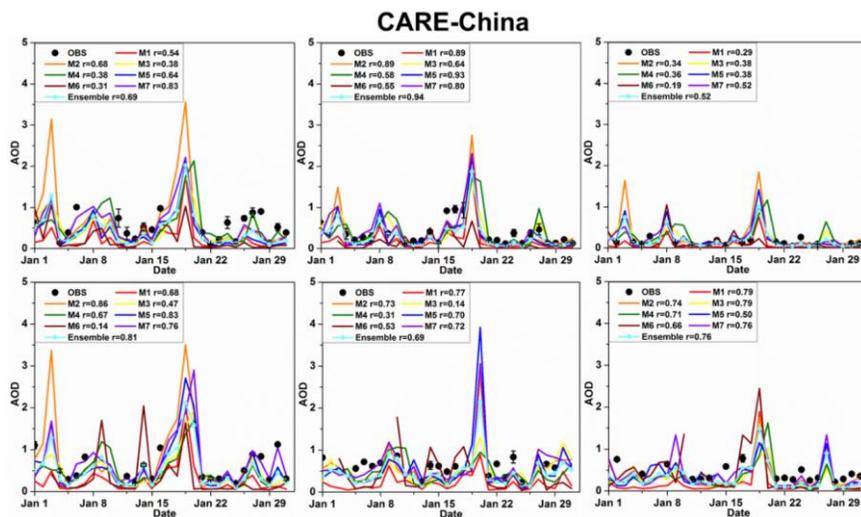


Figure 1412. Comparisons between simulated and observed daily (daytime) mean AOD at the CARE-China sites (Baoding, Beijing City, Beijing Forest, Cangzhou, Jiaozhou, Shenyang,)

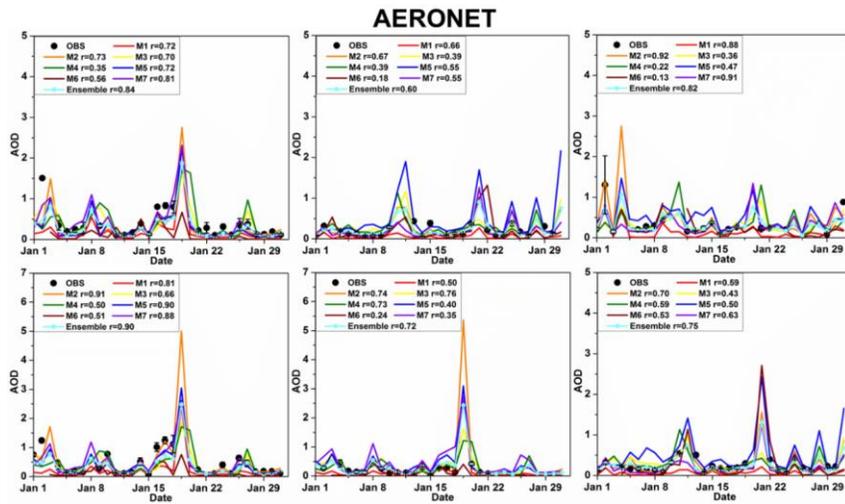


Figure 1413. Comparisons between simulated and observed daily (daytime) mean AOD at the AERONET sites (Beijing, Shirahama, GIST, Xianghe, Xinglong, Osaka)