

We thank the referees for their comments, which helped improving the quality of our manuscript. A point by point response (in black) to the reviewers' comments (in blue) will follow. Changes in the text are indicated in in *black italics*.

#### Reviewer 1:

Organic aerosol (OA) is a major component of the total aerosol. Thus, understanding the sources of OA is crucial for controlling aerosol pollution. In the past decade, the development of the aerosol mass spectrometer allows the direct quantification of organic aerosol mass, and source apportionment can be done by further applying PMF on such data set. The main weakness of AMS is that the hard ionization technique causes significant fragmentation, so that some sources are difficult to separate, such as secondary formation from anthropogenic or biogenic origins. To overcome this drawback, the authors, for the first time, explored the usage of LDI technique on offline samples, which causes less fragmentation and may give new insights of OA sources. The authors developed novel procedures for instrument calibration and uncertainty assessment, generated reference spectra for PMF from both specific environment (in the tunnel for traffic sources) and the lab (for wood burning sources), and carefully evaluated the results by comparing to the results from AMS technique. By doing so, the authors reported both the advantages and some potential issues of this technique, which makes this work easy to understand and follow. Although the ionization efficiency and the "matrix effect" of LDI technique remains poorly understood, which weakened the interpretability of some PMF factors, this work provided several new insights in OA sources and should be published in ACP after the following comments are addressed.

#### Minor comments:

Page 4, Line 16-17. "The observed ions are typically still fragments: : :". Some of the spectra (e.g. Fig.3d) reported in this work contains some high-mass peaks, which might be complete molecules. This work will be strengthened if you can identify them.

We agree with the reviewer that the identification of ions in the recorded mass spectra would greatly help gaining a deeper understanding of sources and processes but also of the measurement technique itself. In the complex mixtures that we analyze, different ions have plausibly a significant contribution to the signal at a single nominal mass, especially given the size of the ions that we observe. With the mass resolution and mass accuracy of the  $m/z$  calibration an identification of the contributions of different ions is not possible unambiguously. Stark et al. (2015) developed an approach to estimate the carbon oxidation state using the mass-to-charge ratio of a peak and its mass defect (difference to closest integer) for mass spectra with limited mass resolution. With our instrumental approach a similar approach cannot be applied because of the poor mass accuracy of the  $m/z$  calibration.

We have adapted the text according to the reviewer's suggestion:

*"... With this accuracy and resolution, neither distinguishing different ions at the same nominal mass nor estimating properties such as O/C and H/C for a nominal mass following Stark et al. (2016) was possible for the quartz-filter-LDI-MS measurements. Instead, the spectra were integrated to UMR sticks. ..."*

As reviewer 2 mentions, the prominent presence of odd masses might hint to a rather prominent fraction of signal attributable to fragments as reviewer 2 mentioned. However, it is currently impossible to distinguish odd mass molecular contributors from fragments. According to Samburova et al. (2005a), fragmentation in this type of instrument is negligible. The mass spectral signatures acquired in this study are similar to the ones in Samburova et al. (2005a) and the lower laser energy applied in this study is lower than in the former study. This suggests that also during our measurements fragmentation was not a prominent process.

We added this discussion to the manuscript:

“ ...

*The presence of considerable signal at odd masses might indicate that significant fragmentation occurs during desorption and ionization of the organic matter in the LDI-MS (Fig. 1a, 3, 4). However, Samburova et al. (2005a) suggested that fragmentation in this instrument is negligible. The mass spectral signatures acquired in this study are similar to the ones in Samburova et al. (2005a) and the lower laser energy applied in this study is lower than in the former study. This suggests that also during our measurements fragmentation was not a prominent process. Overall, the extent of fragmentation remains unclear.*

...”

Page 5, Line 5-20 and Page 7, Line 5-10. The usage of AgNO<sub>3</sub> seems unnecessary to me. First, using UMR data does not require a high-quality mass calibration. Second, you mentioned that AgNO<sub>3</sub> may cause change in ionization physics, which might even hurt the mass calibration. If there are other important reasons for using AgNO<sub>3</sub>, please specify. Otherwise, I think it is important to mention that spiking AgNO<sub>3</sub> should be skipped in future works. Following this comment, the mass calibration accuracy (Fig.3c) is very unstable and too low under the mass resolution of 1000-2000. Do you have explanations?

The m/z calibration parameters in our LDI-MS measurements were highly variable between different samples. Therefore, to determine unit mass resolution integration regions it was necessary to perform an m/z calibration on every single sample.

In other mass spectrometers, some peaks are always observed, independent of the sample, and can therefore be used for an m/z calibration and/or to assess the performance of an initial m/z calibration: E.g. in the Aerodyne AMS the ions N<sub>2</sub><sup>+</sup>, CO<sub>2</sub><sup>+</sup>, W<sup>+</sup>, etc. are present under most conditions. As another example, in chemical ionization mass spectrometry m/z calibration is typically performed using the known chemical reagent ions. For a reliable m/z calibration such known anchor ions are crucial. In many complex mixtures as for ambient aerosol, the signal observed at any nominal mass is composed of ions from numerous organic compounds and peaks are broad. Therefore, the exact mass of such a peak is unknown hindering an accurate m/z calibration.

In matrix assisted LDI-MS matrix molecules or matrix fragment ions have been used previously. However, the native matrix of our samples was highly variable, so that no ions provided suitable anchors for calibration. Therefore, we chose to spike our samples with AgNO<sub>3</sub>. In the absence of such anchor ions an m/z calibration, even to integer accuracy, is not feasible.

Indeed, as mentioned in the reviewer's comment, we observed differences between the average spectrum containing AgNO<sub>3</sub> and the average spectrum without AgNO<sub>3</sub> which we related to changing ionization conditions. Hence, we did not report atmospheric data for AgNO<sub>3</sub>-spiked samples. Rather, we used the AgNO<sub>3</sub>-spiked spectra to obtain a first calibration (based on Ag<sup>+</sup> anchor ions), used the first calibrate to define prominent, filter-specific anchor ions, and then calibrated the unspiked spectra using these anchor ions. This

procedure is explained in Section 2.2.2 Data treatment of the revised manuscript (unchanged from original) and was essential to perform an objective calibration. Because only prominent peaks were used in the second calibration step, the possibility of an influence of  $\text{AgNO}_3$  on these samples is minimal.

Although we had mentioned these issues in the submitted manuscript, we have clarified them further following both reviewers' comments. We adapted the respective part of the manuscript to:

“... ”

*The  $m/z$  calibration parameters were highly variable between different samples. Therefore, to determine unit mass resolution integration regions it was necessary to perform an  $m/z$  calibration on every single sample. However, unlike in the Aerodyne AMS ( $\text{N}_2^+$ ,  $\text{O}_2^+$ ,  $\text{W}^+$ ), there are no dominant anchor ions present in these spectra that could be used for an  $m/z$  calibration. In absence of such ions, we performed a two-step calibration procedure. Each sample was spiked with silver nitrate ( $\text{AgNO}_3$ , Sigma Aldrich, >99.8%) as an internal standard (approach illustrated in Fig. 1). In order to avoid the suppression of the sample signal, the internal standard (aqueous solution 500 ppt to 20 ppm) was only placed (as a droplet) on a small part of the sample. The 499 spectra from all positions on the measurement grid were separated and defined as (1) silver-spiked (lower panel in Fig. 1a), (2) silver-free (upper panel in Fig. 1a), and (3) intermediate-silver. Intermediate-silver cases were defined using the signal intensity in the regions of the mass spectrum where silver was expected, in comparison to adjacent silver-free regions of the spectrum, and were discarded. To calculate the first calibration, we calibrated the average silver-spiked spectrum of each filter sample, using the peaks of the silver monomer ( $m/z$  107, 109 Th), dimer ( $m/z$  214, 216, 218 Th), and trimer ( $m/z$  321, 323, 325, 327 Th). We found that this calibration of silver-spiked was not directly applicable to the average silver-free mass spectrum. Possibly, spiking the filter region with aqueous  $\text{AgNO}_3$  caused enough of a change to the surface of the filter sample to influence the ionization physics. This could affect the  $m/z$  calibration, since the delayed-pulse ion extraction in our instrument is not orthogonal to the ablation plume but nearly parallel. Therefore, a second calibration was obtained for the averaged silver-free spectra using prominent non-silver peaks present in both spectra. Such a two-step calibration is necessary to achieve accurate  $m/z$  calibrations of the silver-free spectra.*

“... ”

The mass resolution as presented in Fig. 1b is computed based on the silver peaks since all other peaks are potentially influenced by a multitude of ions. Mass resolution and accuracy are independent properties; broad peaks can still be accurately calibrated in the absence of interference. We acknowledge that the mass accuracy in our dataset is poor. Therefore, we emphasize in the text that based on such accuracy an identification of contributing ions to a nominal mass is not possible (i.e. high-resolution analysis).

*“... With this accuracy and resolution, neither distinguishing different ions at the same nominal mass nor estimating properties such as O/C and H/C for a nominal mass following Stark et al. (2016) was possible for the quartz-filter-LDI-MS measurements. Instead, the spectra were integrated to UMR sticks...”*

Since the ions are not orthogonally extracted into the ToF unit, variability in the ionization process might be propagated into the ToF unit and, thereby, contribute to the poor accuracy of the  $m/z$  calibration. Since this variability was observed for the 1<sup>st</sup> calibration, which used only the Ag “anchor” ions which dominated the spectral intensity (so that a substantial interference by unidentified ions is very unlikely), it must be generally attributed to variability in the ionization and detection processes.

An observation of calibration inaccuracy in the calibration itself indicates that the calibration model was imprecise. The calibration model is the standard equation describing ion time-of-flight in the mass spectrometer, and would be violated if ions were at different distances from the detector or had different kinetic energy vectors, upon extraction. A non-uniform electric field in the spectrometer could also cause aberrances. These problems have been minimized in some instruments, but their minimization would reduce our ability to directly analyze filter samples, which is a major strength of our approach.

Page 5, Line 31. Remove “scales”

We removed this word.

Page 5, Line 34. It is the first time you mention “WSOC”, please mention the full name.

We added this information:

*“...  $OC_{Sunset}$  was determined using the Sunset OC/EC analyzer and  $(WSOM/WSOC)_{oAMS}$  through WSOM measurements (described in Section 2.1, WSOC being water-soluble organic carbon). ...”*

Page5, Eq1. Scaling LDI intensity to AMS intensity may cause problems. In later analyses, we know that LDI detection has bias on different OA sources, but AMS detection should be rather equal. If the fraction of different OA sources changed, for example, LDI-insensitive source increased but LDI-sensitive source decreased in the same magnitude, this change will be clearly seen as a decrease in LDI data, but NOT in AMS data. Under such circumstances, the scaling factor has a dependency on source distribution. This might (NOT necessarily) affect the overall PMF results, for example the seasonal variation of PMF factors and the comparison of LDI-PMF and AMS-PMF. Thus, I think it is necessary to double check the results using unscaled LDI data.

The reviewer addresses one of the main challenges related to LDI-MS analyses: the quantification. During our experiments we discovered that (1) the signal recorded for the same sample during 1 day is fairly variable introducing a lot of variability and (2) that instrumental changes over time (possibly properties of the laser) introduced a decrease of signal with the measurement time. Consequently, the raw signal does not show a relation with OC, the sum of OC and EC or PM10 and is, therefore, not interpretable (Fig. S3, S4). Therefore, scaling to an external property is inevitable.

We updated the offline AMS source apportionment results used in the comparison (Fig 7) with the recently published data in Daellenbach et al. (2017). In the same time we also reworked Fig. 9 and 10 for an improved visibility. This resulted in small changes in Fig. 7 and the related text:

“ ...

*As was the case for the comparison to  $NO_x$  and  $eBC_{tr}$  (Section 3.4.2), also in the comparison to the offline AMS traffic (HOA), a higher LDI-traffic/HOA was observed in summer which contributes to the low correlation coefficient ( $R^2=0.04$ ). This might be related to the fact that HOA is only primary*

*without major contributions of traffic SOA or dust resuspension while LDI-traffic is potentially also influenced by aged/secondary traffic aerosol and resuspension. In contrast to summer, we observe a good agreement between LDI-traffic and HOA in winter. Thus even though a varying relative response factor (rRF) of the LDI-MS might contribute to these differences, these biases are not systematic but season dependent. As stated earlier, LDI-traffic is thought to be a mixture of primary tailpipe exhaust, aged/secondary tailpipe exhaust, and resuspended dust (as well as tyre break and engine wear). This will be further elucidated in Section 3.4.5.*

*LDI-BB was highly correlated with offline AMS-BBOA ( $R^2=0.83$ ), yet the LDI-BB concentrations were higher than BBOA from oAMS, especially in northern Switzerland (with an LDI-BB:AMS-BBOA ratio between 1.4 and 4.1 for the different sites). A possible reason is also here the mixing of secondary components into LDI-BB when comparing to primary BBOA from oAMS. However, we cannot exclude that this effect is due to different rRFs of the LDI-MS for different compound classes.*

*The identified secondary components, LMW-OA ( $R^2=0.45$ ) and bio-OA ( $R^2=0.62$ ), also correlate with the corresponding OOA factors from the oAMS analysis, WOOA and SOOA, yet the correlation coefficients are smaller than for LDI-BB and BBOA. For these factors, differences between the two methodologies could be related to differences in the PMF performance or to differences in the response factors for different components in the LDI-MS (the LMW-OA: WOOA ratio is between 0.7 and 1.1 and the bio-OA: SOOA ratio between 0.2 and 0.4).*

*...*

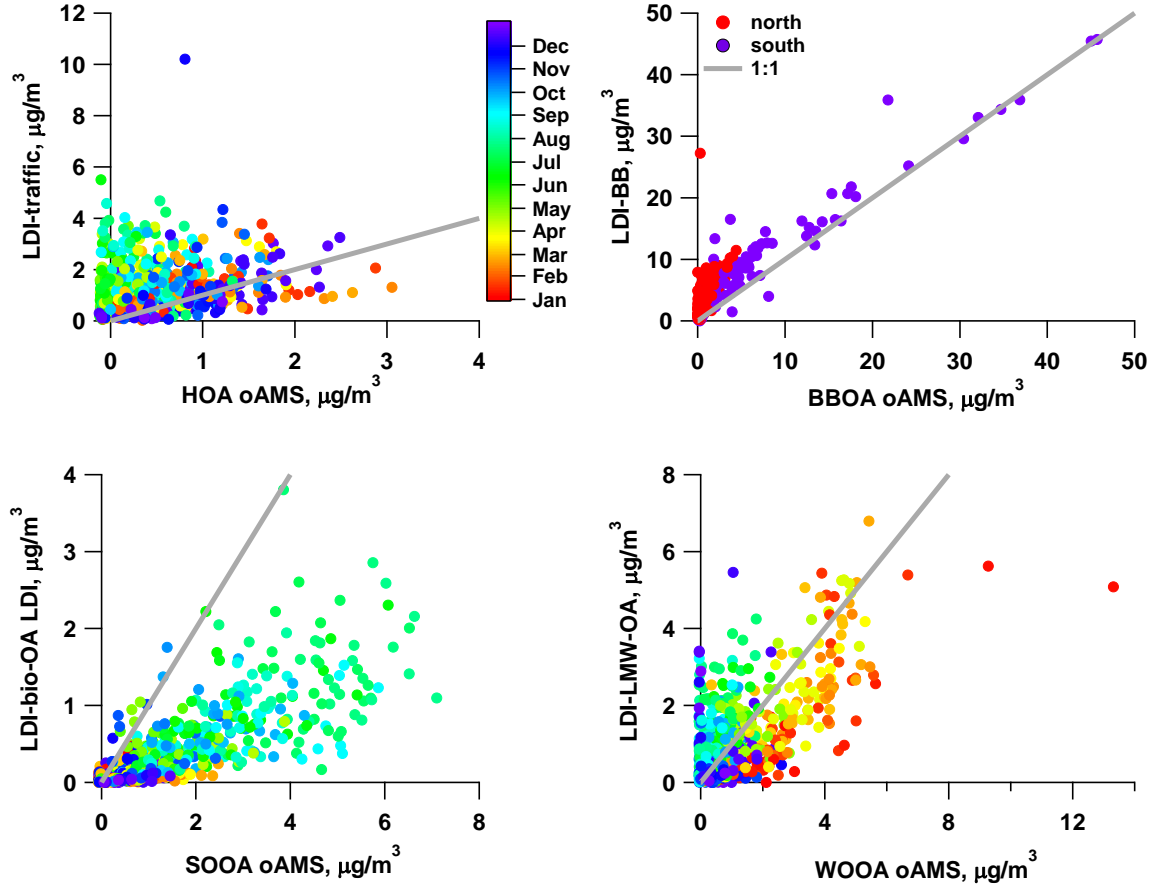


Figure 1: Comparison of LDI-MS to reference offline AMS source apportionment results for the sum of traffic related factors (LDI-traffic,  $R^2_{\text{LDI,oAMS}}=0.04$ ), sum of wood burning related factors (LDI-BB,  $R^2_{\text{LDI,oAMS}}=0.83$ ), bio-OA ( $R^2_{\text{LDI,oAMS}}=0.62$ ), and LMW-OA ( $R^2_{\text{LDI,oAMS}}=0.45$ ).

In response to the question, we added a paragraph on the analysis of relative response factors for LDI-MS in Section 3.4.3 of the manuscript:

“ ...

Overall, the source apportionment results based on LDI-MS data provide source separations with similar temporal behaviors as offline AMS and online AMS and ACSM analyses, yet seem to overestimate combustion related primary particle sources and underestimate secondary OA. However, we assume for this comparison that all factors give an equal response at a given concentration, i.e. the relative response factor ( $r\text{RF}_{\text{LDI}}$ ) of all factors to be 1. The relative contribution of a factor  $k$  to the total signal observed with the LDI-MS ( $r_{i,k,\text{LDI}}$ ) a specific measurement ( $i$ ) depends on the factor concentration ( $g_{i,k,\text{LDI}}$ ) as well as the sum of all factors ( $\sum_k^p(g_{i,k,\text{LDI}})$ ) separated for the LDI-MS data. Assuming that differences between the LDI-PMF and the AMS-PMF arise solely from different  $r\text{RF}_{\text{LDI}}$ s for different factors,  $r_{i,k,\text{LDI}}$  is a function of the AMS factor concentrations ( $g_{i,k,\text{AMS}}$ ) and the  $r\text{RF}_{\text{LDI}}$  of this factor:

$$r_{i,k,LDI} = \frac{g_{i,k,LDI}}{\sum_k^p(g_{i,k,LDI})} = \frac{g_{i,k,AMS} * rRF_{k,LDI}}{\sum_k^p(g_{i,k,AMS} * rRF_{k,LDI})} \quad (4)$$

In order to determine  $rRF_{LDI}$  for the LDI-MS several strong assumptions are required: (1) the sum of traffic1 and traffic2 represents HOA, (2) the sum of WBeff, WBineff1 and WBineff2 represents BBOA, (3) LMW-OA represents WOOA, (4) bio-OA represents SOOA and (5) the AMS factors for which there is an LDI-MS equivalent, (1) - (4), are the only contributors to OA. In scenario 1 All the above assumptions are considered to be true. For scenario 1, we estimate  $rRF_{LDI-BB}$  as 1.84,  $rRF_{LDI-LMW-OA}$  as 0.29 and  $rRF_{LDI-bio-OA}$  as 0.40 (using LDI-traffic as reference factor,  $rRF_{LDI-traffic} = 1.00$ ). The LDI-MS factor concentrations corrected using  $rRF_{LDI}$  show a close relation to the uncorrected LDI factor concentrations ( $R^2_{LDI-traffic}=0.78$ ,  $R^2_{LDI-BB}=0.94$ ,  $R^2_{LDI-LMW-OA}=0.80$ ,  $R^2_{LDI-bio-OA}=0.85$ , Fig. 8).

We tested the sensitivity of the  $rRF_{LDI}$  estimates to the assumptions (1) and (2) being wrong (scenario 2) and only assumption (5) being wrong (scenario 3). In scenario 2, we alter scenario 1 by comparing the sum of traffic1 and bio-OA to SOOA and the sum of BBineff2 and LMW-OA to WOOA when computing  $rRF_{LDI}$ . For scenario 2, LDI-MS factor concentrations corrected using  $rRF_{LDI}$  show also a close relation to uncorrected LDI-MS factor concentrations ( $R^2_{LDI-traffic2}=0.96$ ,  $R^2_{BBeff+BBineff1}=1.00$ ,  $R^2_{LDI-BBineff2+LMW-OA}=0.99$ ,  $R^2_{traffic1+bio-OA}=0.96$ ). In scenario 3, we alter scenario 1 by considering also AMS factors without an equivalent in the LDI-MS PMF. Therefore, we compare LDI-traffic to the sum of HOA, COA, and SC-OA. Under these conditions LDI-MS factor concentrations corrected using  $rRF_{LDI}$  show a close relation to uncorrected LDI-MS factor concentrations ( $R^2_{LDI-traffic}=0.75$ ,  $R^2_{LDI-BB}=0.90$ ,  $R^2_{LMW-OA}=0.74$ ,  $R^2_{bio-OA}=0.85$ ). The differences in  $rRF_{LDI}$  between scenarios 1, 2, and 3 highlight the uncertainties caused by plausible violations of the underlying assumptions when determining  $rRF_{LDI}$ . Overall the estimates  $rRF_{LDI}$  are highly uncertain and their accurate determination needs to be the focus of future work. Given the high uncertainty of  $rRF_{LDI}$  and the good correlation between  $rRF_{LDI}$  corrected and uncorrected LDI-MS factor concentrations, we present uncorrected results without considering  $rRF_{LDI}$ .

**Table 1: relative response factors (rRF) for LDI-MS analyses for 3 different scenarios.**

scenario 1	scenario 2	scenario 3
$rRF_{LDI-traffic} = 1.00 \pm 0.00$	$rRF_{LDI-traffic2} = 1.00 \pm 0.00$	$rRF_{LDI-traffic} = 1.00 \pm 0.00$
$rRF_{LDI-BB} = 1.84 \pm 0.06$	$rRF_{LDI-BBeff+LDI-BBineff1} = 1.26 \pm 0.06$	$rRF_{LDI-BB} = 5.67 \pm 0.13$
$rRF_{LDI-LMW-OA} = 0.29 \pm 0.02$	$rRF_{LDI-LMW-OA+LDI-BBineff2} = 1.67 \pm 0.07$	$rRF_{LDI-LMW-OA} = 0.30 \pm 0.02$
$rRF_{LDI-bio-OA} = 0.14 \pm 0.01$	$rRF_{LDI-bio-OA+LDI-traffic1} = 0.39 \pm 0.02$	$rRF_{LDI-bio-OA} = 0.65 \pm 0.02$

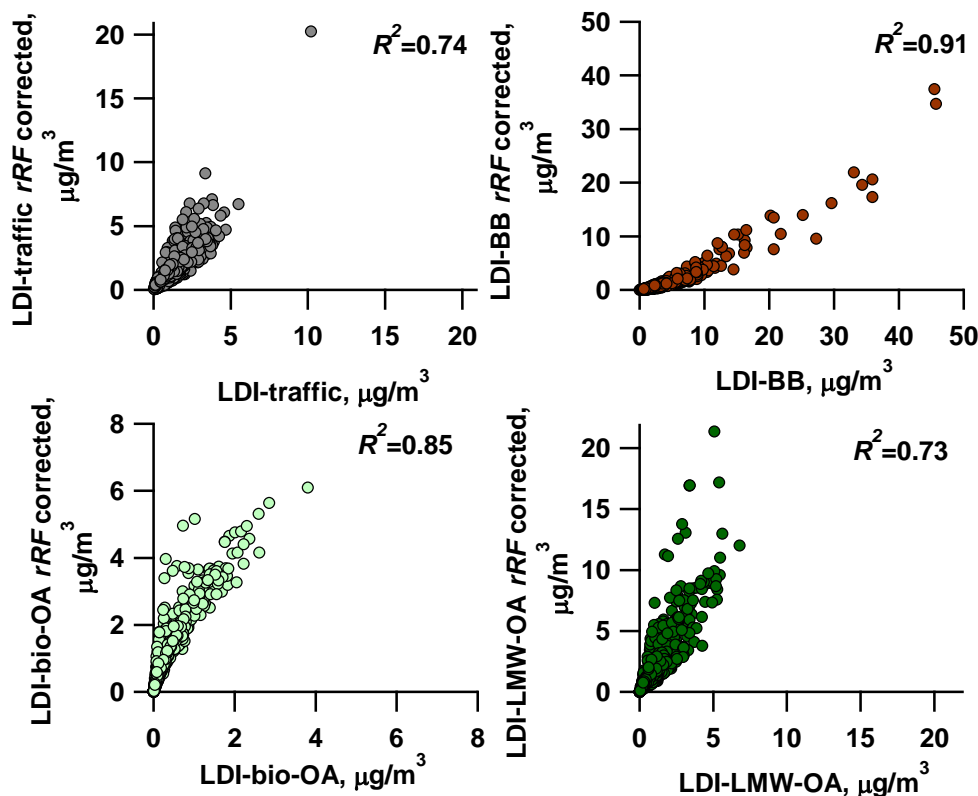


Figure 2: Comparison of LDI-MS factor concentrations corrected using relative response factors ( $rRF$ ) to uncorrected LDI factor concentrations (Scenario 1).

...”

Additionally, we also add a comment in the conclusion:

“... Other components are more difficult to compare quantitatively because of different source separations in PMF as well as differences in the relative response factors ( $rRF$ ) of OA components.  $rRF$  determined in this study are uncertain and, therefore, not used for correcting the LDI-PMF results. ...”

Page 7, Line 22-25. The fitting parameter in November is clearly different (Fig.2 red dots), though maybe not significantly.

The purpose of these exercises was to obtain an error model as presented in Section 3.1 and in the Supplement. There are only three repeats. Therefore, it is hard to judge whether the experiment in question is an outlier or the result of a broad distribution. We report the resulting uncertainty of the error model in the manuscript as following:

“... The relative error was related to the absolute signal intensity for each  $m/z$  for each sample and experiment (Fig. 2), with an asymptotic value of only  $\sim 9 \pm 7 \%$  at high signals (average and standard deviation of fitting Eq. 3). The average absolute error for small signals was  $102 \pm 48$  a.u. (from Eq. 3)...”



Page 8, Line 33. “: : of which 84, 94, 120, and 177 correlated with NO<sub>x</sub>...”. 84amu is not marked in Fig.4a, and it does not show a strong correlation with NO<sub>x</sub> (R ~0.2 as I read from the figure)

The sentence has been corrected:

*“... The m/z 84 94, 101, 120, 143, 165, and 177 Th, also detected in tunnel samples, showed high relative intensities in Bern in summer, of which 94, 120, and 177 Th correlated with NO<sub>x</sub>, which suggests a high contribution of traffic emissions at this location, consistent with previous observations (Zotter et al., 2014). ...”*

Page 9, Line 15. “Both traffic1 and traffic2 showed a relationship to NO<sub>x</sub>, only traffic2 correlated with eBC<sub>tr</sub> (Table S1)”. What does the “relationship” mean? The correlation between traffic 1 and NO<sub>x</sub> is very weak. From Table S1, I think the traffic 1 has neither correlation with NO<sub>x</sub>, nor with eBC; while traffic 2 may correlate with both.

We rephrased the respective part of the manuscript:

*“... Both traffic1 and traffic2 showed increasing concentrations with increasing NO<sub>x</sub> levels. However traffic1/NO<sub>x</sub> and traffic2/NO<sub>x</sub> were seasonally variable ( $R_{p,traffic1,NO_x}=-0.17$ ,  $R_{p,traffic2,NO_x}=0.24$ , Fig. 6, Fig. S9). While traffic2 correlated with eBC<sub>tr</sub>, traffic1 did not show such a dependency (Tab. S1, Fig. 6, Fig. S9). ...”*

Page 9, Line 16-17. “This suggests that traffic2 is related to primary emissions from the combustion process and that traffic2 is also influenced by other processes, like e.g. secondary production.” For people who are not familiar with traffic tracers, e.g. NO<sub>x</sub> and eBC, this interpretation is hard to follow. It is good to explain this a bit more.

We rephrased the respective part of the manuscript:

*“... The correlation with eBC<sub>tr</sub> suggests that traffic2 is related to primary emissions from the combustion process. Based on the lack of correlation between traffic1 and eBC<sub>tr</sub>, traffic1 might also be influenced by other processes, such as e.g. secondary production. ...”*

Page 10, Line 31-33. “The contribution of the last remaining factor showed an exponential increase with temperature, similar to terpene emissions and biogenic SOA (Leaith et al., 2011), suggesting this factor to be strongly influenced by biogenic SOA production. Therefore, this factor was termed bio-OA.” Some other atmospheric processes, other than VOC emissions, are temperature-dependent, such as VOC oxidation and photochemistry. Is there additional evidence to validate the bio-OA factor, such as higher fractional contribution in rural sites than in urban sites?

Using yearly average NO<sub>x</sub> concentrations as a metric for rurality/urbanity, Payerne sticks out as rural (low NO<sub>x</sub>) and Bern as urban (high NO<sub>x</sub>). In Payerne bio-OA has a higher relative contribution (11%) than in Bern (3%) which are the highest and lowest in the entire dataset (see Table 2 in manuscript). Qualitatively, the same was observed in the AMS data (38% SOOA in Payerne and 21%; only S. Vittore showed a lower contribution at the expense of higher BBOA, Daellenbach et al., 2017).

In order to make this information more visible we added yearly NO<sub>x</sub> concentrations to Tab. 2 and sorted the sites according to the yearly average NO<sub>x</sub> concentration.

Such information was added to the manuscript:

*“...Bio-OA had a highest relative contribution at the most rural site in the dataset (Payerne: yearly average bio-OA 11%, yearly average NO<sub>x</sub> concentration: 9 ppb) and the lowest at the most trafficated site (Zurich: yearly average bio-OA 3%, yearly average NO<sub>x</sub> concentration: 48 ppb). ...”*

#### **References:**

Daellenbach, K. R., Stefenelli, G., Bozzetti, C., Vlachou, A., Fermo, P., Gonzalez, R., Piazzalunga, A., Colombi, C., Canonaco, F., Hueglin, C., Kasper-Giebl, A., Jaffrezo, J.-L., Bianchi, F., Slowik, J. G., Baltensperger, U., El-Haddad, I., and Prévôt, A. S. H.: Long-term chemical analysis and organic aerosol source apportionment at nine sites in central Europe: source identification and uncertainty assessment, *Atmos. Chem. Phys.*, 17, 13265-13282, <https://doi.org/10.5194/acp-17-13265-2017>, 2017.

We thank Reviewer 2 for their comments, which helped improving the quality of our manuscript. A point by point response (in black) to the reviewers' comments (in blue) will follow. Changes in the text are indicated in in *black italics*.

#### Reviewer 2:

Mass spectrometrical methods are in widespread use to characterize the organic composition of particulate matter. Aerosol mass spectrometers, despite their drawbacks with respect to hard ionization, are often drawn upon for source apportionment studies in combination with PMF. Here, the authors applied a LDI technique for an extensive study of ambient aerosol samples. Measurements and data analysis have been conducted meticulously, also reflecting on uncertainties, and source apportionment results have been compared to AMS data. The work should be published in ACP after some minor revisions.

1) There is not much detail on the experimental setup of the mass spectrometrical experiment. Some hints are given, such as "ion extraction is not orthogonal to the ablation plume", then how is it performed?, or information about mass resolution. In the introduction, the adjustment of laser power is mentioned as one of the advantages of LDI-MS compared to ATOF-MS, but information on the laser power in this experiment is missing save the fact that it is adjustable. Therefore, the authors are encouraged to give more details on the laser/mass spectrometer system.

In response to this question, we added the further information on the laser power and ion extraction to the following paragraphs and adapted the sentence on the comparison between ATOFMS and LDI-MS:

*"... Possibly, spiking the filter region with aqueous AgNO<sub>3</sub> caused enough of a change to the surface of the filter sample to influence the ionization physics. This could affect the m/z calibration, since ion extraction in our instrument is not orthogonal to the ablation plume but a delayed pulsed extraction. ..."*

*"...*

*We recorded the mass spectra of 819 filter samples in an m/z range 65-500 thomsons (Th; 1 Th = 1 Da e<sup>-1</sup>, where e is the elementary charge, ion gate at m/z 60 Th) using a laser-desorption/ionization-ToF MS (Shimadzu Axima Confidence, Shimadzu-Biotech Corp., Kyoto, Japan) equipped with an N<sub>2</sub> laser (wavelength 337 nm, frequency 50 Hz, laser pulse width 3 ns, 130-180 µJ/pulse) in the positive reflectron mode. All the accessible instrumental parameters were kept constant during the whole period of measurements taking place from November 2015 to mid-March 2016. Specifically, the laser intensity was adjustable by means of a rotating wheel of filters with varying transmissivity (0 being blocked and 180 being completely open). We set the wheel parameter to 105 of 180 which would result in an estimated laser energy of ~6-9 µJ/pulse or 2.8-4.2\*10<sup>8</sup> W/cm<sup>2</sup> with a 3ns pulse and 30µm laser beam diameter. While the laser energy was initially set and kept constant, the aging of the laser during the given time period was also expected to reduce its intensity. We monitored and assessed changes in laser power and other instrumental parameters, as well as possible sources of uncertainty/contamination from sample preparation and intra and inter-day reproducibility, by repeated measurements of a subset of our samples.*

*..."*

*"... Similar to the AMS, online single-particle LDI-MS instruments such as the ATOFMS also yield extensive fragmentation. However, such fragmentation can be avoided by measuring offline aerosol samples (filters) using other systems. Samburova et al. (2005a) showed by comparing*

*measurement with and without matrix addition, that fragmentation was negligible in their instrument (wavelength 337 nm, LDI-MS, Shimadzu/Kratos, Axima CFR). ...”*

2) The mass resolution is low in this experiment, so it is understandable that the authors did not assign their signals to actual compounds. But I wonder, is there not the possibility to assign at least some prominent peaks relating to past experience or literature data? In this respect, there are many odd mass numbers in the spectra, hinting at the formation of fragment ions. This should be discussed briefly.

As the reviewer points out it is impossible to distinguish the contribution of different ions at the same nominal mass. We agree with the reviewer that it is desirable to assign compounds to the most prominent masses and that the prominent presence of odd masses hints to a rather prominent fraction of signal attributable to fragments. However, it is currently impossible to distinguish odd mass molecular contributors from fragments.

In absence of vast literature databases with the same instrument, we prefer to refrain from interpreting the fingerprints further than by comparing to reference samples since we fear to over-interpret the data. Earlier work confirms the prevalence of the same peaks for 1 site present in our dataset but cannot attribute these peaks to specific compounds (Zurich, Kaserne, Samburova et al., 2005).

As reviewer 2 mentions, the prominent presence of odd masses might hint to a rather prominent fraction of signal attributable to fragments as reviewer 2 mentioned. However, it is currently impossible to distinguish odd mass molecular contributors from fragments. According to Samburova et al. (2005a), fragmentation in this type of instrument is negligible. The mass spectral signatures acquired in this study are similar to the ones in Samburova et al. (2005a) and the lower laser energy applied in this study is lower than in the former study. This suggests that also during our measurements fragmentation was not a prominent process. We mention the possible influence of fragmentation now also in the main text:

“ ...

*The operational mass resolution,  $m/\Delta m$ , based on the measurements conducted on one sample tray per month and the accuracy of both calibration steps are shown in Fig. 1b and 1c, respectively. The operational resolution of the instrument was only determined for the silver mono- (resolution 1100), di- (1700), and trimer (2100) as we are confident of the absence of interfering ions for these peaks. In comparison to the LDI-MS, the HR-ToF AMS in V-mode (and W-mode) has a resolution of ca. 2'000 (4'000) at  $m/z$  100 Th and an  $m/z$  calibration accuracy < 20 ppm (< 10 ppm, DeCarlo et al., 2006). With this accuracy and resolution, neither distinguishing different ions at the same nominal mass nor estimating properties such as O/C and H/C for a nominal mass following Stark et al. (2016) was possible for the quartz-filter-LDI-MS measurements. Instead, the spectra were integrated to UMR sticks. The presence of considerable signal at odd masses might indicate that significant fragmentation occurs during desorption and ionization of the organic matter in the LDI-MS (Fig. 1a, 3, 4). However, Samburova et al. (2005a) suggest that fragmentation in this instrument is negligible. The mass spectral signatures acquired in this study are similar to the ones in Samburova et al. (2005a) and the lower laser energy applied in this study is lower than in the former study. This suggests that also during our measurements fragmentation was not a prominent process. Overall, the extent of fragmentation remains unclear.*

...”

3) Figure 3 reveals, if there are a lot of signals resulting from a sample, a big unresolved hump is visible ( as can be seen from the wood burning in comparison to the clean, well

structured spectrum from tunnel weekend). Can this be improved/targeted by reducing the laser power, or will this lead to a severe and unacceptable loss in sensitivity?

For such separations, we suggest to extract PM with different solvents in order to get a better idea on their properties in future work. Samburova et al. (2005) observed such humps also in water-soluble PM for 1 of the sites present also in this study (Zurich, Kaserne).

4) What is the gain in spiking with silver nitrate? In the end, you have to calibrate the spectra without showing peaks from silver ions with an independent calibration procedure, as is mentioned in the manuscript. Why then not apply the second procedure, for which obviously no spiking is necessary, in general and omit the silver nitrate solution?

Both reviewers mentioned a similar point. We have taken this feedback as an opportunity to better clarify our two-step calibration procedure, as detailed more fully in the following (the same response has been written to both reviewers):

The  $m/z$  calibration parameters in our LDI-MS measurements were highly variable between different samples. Therefore, to determine unit mass resolution integration regions it was necessary to perform an  $m/z$  calibration on every single sample.

In other mass spectrometers, some peaks are always observed, independent of the sample, and can therefore be used for an  $m/z$  calibration and/or to assess the performance of an initial  $m/z$  calibration: E.g. in the Aerodyne AMS the ions  $N_2^+$ ,  $CO_2^+$ ,  $W^+$ , etc. are present under most conditions. As another example, In chemical ionization mass spectrometry  $m/z$  calibration is typically performed using the known chemical reagent ions. For a reliable  $m/z$  calibration such known anchor ions are crucial. In many complex mixtures as for ambient aerosol, the signal observed at any nominal mass is composed of ions from numerous organic compounds and peaks are broad. Therefore, the exact mass of such a peak is unknown hindering an accurate  $m/z$  calibration.

In matrix assisted LDI-MS matrix molecules or matrix fragment ions have been used previously. However, the native matrix of our samples was highly variable, so that no ions provided suitable anchors for calibration. Therefore, we chose to spike our samples with  $AgNO_3$ . In the absence of such anchor ions an  $m/z$  calibration, even to integer accuracy, is not feasible.

Indeed, as mentioned in the reviewer's comment, we observed differences between the average spectrum containing  $AgNO_3$  and the average spectrum without  $AgNO_3$  which we related to changing ionization conditions. Hence, we did not report atmospheric data for  $AgNO_3$ -spiked samples. Rather, we used the  $AgNO_3$ -spiked spectra to obtain a first calibration (based on  $Ag^+$  anchor ions), used the first calibrate to define prominent, filter-specific anchor ions, then calibrated the unspiked spectra using these anchor ions. This procedure is explained in Section 2.2.2 Data treatment of the revised manuscript (unchanged from original) and was essential to enabling an objective calibration to be performed. Because only prominent peaks were used in the second calibration step, the possibility of an influence of  $AgNO_3$  on these samples is minimal.

Although we had mentioned these issues in the submitted manuscript, we have clarified them further following both reviewers' comments. We adapted the respective part of the manuscript to:

“ ...

*The  $m/z$  calibration parameters were highly variable between different samples. Therefore, to determine unit mass resolution integration regions it was necessary to perform an  $m/z$  calibration on every single sample. However, unlike in the Aerodyne AMS ( $N_2^+$ ,  $O_2^+$ ,  $W^+$ ), there are no dominant*

anchor ions present in these spectra that could be used for an  $m/z$  calibration. In absence of such ions, we performed a two-step calibration procedure. Each sample was spiked with silver nitrate ( $\text{AgNO}_3$ , Sigma Aldrich, >99.8%) as an internal standard (approach illustrated in Fig. 1). In order to avoid the suppression of the sample signal, the internal standard (aqueous solution 500 ppt to 20 ppm) was only placed (as a droplet) on a small part of the sample. The 499 spectra from all positions on the measurement grid were separated and defined as (1) silver-spiked (lower panel in Fig. 1a), (2) silver-free (upper panel in Fig. 1a), and (3) intermediate-silver. Intermediate-silver cases were defined using the signal intensity in the regions of the mass spectrum where silver was expected, in comparison to adjacent silver-free regions of the spectrum, and were discarded. To calculate the first calibration, we calibrated the average silver-spiked spectrum of each filter sample, using the peaks of the silver monomer ( $m/z$  107, 109 Th), dimer ( $m/z$  214, 216, 218 Th), and trimer ( $m/z$  321, 323, 325, 327 Th). We found that this calibration of silver-spiked was not directly applicable to the average silver-free mass spectrum. Possibly, spiking the filter region with aqueous  $\text{AgNO}_3$  caused enough of a change to the surface of the filter sample to influence the ionization physics. This could affect the  $m/z$  calibration, since the delayed-pulse ion extraction in our instrument is not orthogonal to the ablation plume but nearly parallel. Therefore, a second calibration was obtained for the averaged silver-free spectra using prominent non-silver peaks present in both spectra. Such a two-step calibration is necessary to achieve accurate  $m/z$  calibrations of the silver-free spectra.

..."

#### 5) What is the difference between WSOM and WSOC?

WSOC refers to the water-soluble organic carbon concentration only considering the carbon mass, i.e. number carbon atoms multiplied with the mass weight of carbon per  $\text{m}^3$ . WSOM considers also all other atoms present in the organic molecules (e.g., oxygen, hydrogen, nitrogen, sulfur) by determining a ratio organic matter per organic carbon (OM/OC).

We completed the manuscript with this information:

"... This involved the analysis of the water-soluble organic matter (WSOM) by a high resolution time-of-flight AMS. ..."

"...  $\text{OC}_{\text{Sunset}}$  was determined using the Sunset OC/EC analyzer and  $(\text{WSOM}/\text{WSOC})_{\text{oAMS}}$  through WSOM measurements (described in Section 2.1, WSOC being water-soluble organic carbon). ..."

#### 6) Page 2, line 2: Omit the impact with EI, call it just electron ionization.

We corrected the manuscript.

"... Instruments equipped with electron ionization (EI), such as the aerosol mass spectrometer (AMS, Canagaratna et al., 2007) and aerosol chemical speciation monitor (ACSM, both Aerodyne Research, Inc., Ng et al., 2011a; Fröhlich et al., 2013) provide quantitative online measurements of OA (Jimenez et al., 2016). ..."

#### 7) Please use "such as" instead of "like" in elaborating insertions.

We corrected the respective parts of the manuscript.

"... Although the importance of matrix effects is evident, studies with a single-particle LDI-MS (ATOFMS, laser wavelength 266 nm) have achieved good correlations of major aerosol components with well-established reference measurements such as elemental carbon (EC, Sunset Analyzer, Sunset Laboratory Inc.), OA (AMS),  $\text{NH}_4$  (AMS),  $\text{SO}_4$  (AMS),  $\text{NO}_3$  (AMS), and K (collected



*with a particle-into-liquid sampler, PILS), suggesting that underlying matrix effects did not dominate measurement reproducibility (Healy et al., 2013). ..."*

*"...Similar to the AMS, online single-particle LDI-MS instruments such as the ATOFMS also yield extensive fragmentation.. However, such fragmentation can be avoided by measuring offline aerosol samples (filters) using other systems. ..."*

*"... The correlation with  $eBC_{tr}$  suggests that traffic2 is related to primary emissions from the combustion process. Based on the lack of correlation between traffic1 and  $eBC_{tr}$ , traffic1 might also be influenced by other processes, such as e.g. secondary production. Additionally, tailpipe exhaust cannot be distinguished from other traffic related emissions in the tunnel samples and might also contribute to traffic1. ..."*

*"... The north-south gradient in the  $BB/K^+$  ratios might be caused by imperfections of the separation of the burning conditions by PMF, but other effects such as the age of the stove population and the used technology could contribute as well. ..."*

8) Page 8, line 28: Please do not use Thomson as unit of m/z differences. Maybe Dalton or stay with m/z.

We prefer to use Thomson as a unit for m/z for the case that some ions have more than one charge. We now include a statement on the unit of m/z that we use in this manuscript, including a description on how Thomson (Th) relate to Dalton ( $Th = Da \cdot e^{-1}$ ). Additionally, we also added the unit of m/z throughout the manuscript wherever it was missing.

*"... We recorded the mass spectra of 819 filter samples in an m/z range 65-500 thomsons ( $Th$ ;  $1 Th = 1 Da \cdot e^{-1}$ , where  $e$  is the elementary charge, ion gate at m/z 60  $Th$ ) using a laser-desorption/ionization-ToF MS (Shimadzu Axima Confidence, Shimadzu-Biotech Corp., Kyoto, Japan) equipped with an N2 laser (wavelength 337 nm, frequency 50 Hz, laser pulse width 3 ns, 130-180  $\mu J$ /pulse) in the positive reflectron mode. ..."*

*"... In general, the mass spectra from winter-time were characterized by higher contributions from higher-molecular-weight fragments compared to summer. For the winter sample from San Vittore, a large "hump" was present, on top of which signals appeared with high intensities and a highly regular pattern of m/z differences of 14  $Th$ . ..."*

9) Page 13, line 6: In the first sentence of the conclusion, I would not call it a novel method, rather the utilization of a known method in use since approx. 2000 with some new developed novel aspects.

We changed the manuscript according to the reviewer's suggestion.

*"... In this study, we advanced a known method for the chemical characterization of particulate matter collected on quartz-fiber filters by LDI-MS and applied the method to 819 samples. ..."*

## References:

Samburova, V., Szidat, Hueglin, C., Fisseha, R., Baltensperger, U., Zenobi, R., and Kalberer, M.: Seasonal variation of high-molecular-weight compounds in the water-soluble fraction of organic urban aerosols, *J. Geo. Res.*, 110, D23, doi:10.1029/2005jd005910, 2005.

Murray, K. K., Boyd, R. K., Eberlin, M. N., Langley, J. G., Liang, L., and Naito, Y.: Definitions of terms relating to mass spectrometry (IUPAC Recommendations 2013), *Pure Appl. Chem.*, 85, 1515-1609, doi:10.1351/PAC-REC-06-04-06, 2013.

# Insights into organic-aerosol sources via a novel laser-desorption/ionization mass spectrometry technique applied to one year of PM<sub>10</sub> samples from nine sites in central Europe

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**Abstract.** We assess the benefits of offline laser-desorption/ionization mass spectrometry (LDI-MS) in understanding ambient particulate matter (PM) sources. The technique was optimized for measuring PM collected on quartz-fiber filters using silver nitrate as an internal standard for  $m/z$  calibration. This is the first application of this technique to samples collected at nine sites in central Europe throughout the entire year 2013 (819 samples). Different PM sources were identified by positive matrix factorization (PMF) including also concomitant measurements (such as NO<sub>x</sub>, levoglucosan, and temperature). By comparison to reference mass spectral signatures from laboratory wood burning experiments as well as samples from a traffic tunnel, three biomass-burning factors and two traffic factors were identified. The wood-burning factors could be linked to the burning conditions; the factors related to inefficient burns had a larger impact on air quality in southern Alpine valleys than in northern Switzerland. The traffic factors were identified as primary tailpipe exhaust and most possibly aged/secondary traffic emissions, respectively. The latter attribution was supported by radiocarbon analyses of both the organic and elemental carbon. Besides these sources, also factors related to secondary organic aerosol were separated. The contribution of the wood burning emissions based on LDI-PMF correlates well with that based on AMS-PMF analyses, while the comparison between the two techniques for other components is more complex.

## 1 Introduction

Climate and health are strongly affected by atmospheric aerosols (Kelly et al., 2007; IPCC, 2013), a substantial fraction of which is organic (Jimenez et al., 2009 and reference therein). This organic aerosol (OA) is a complex mixture of thousands of compounds (Goldstein and Galbally, 2007), of which only 10-30% have been speciated by modern techniques (Hoffmann et al., 2011; Simoneit et al., 2004). Therefore, the chemical composition of OA, its emission sources and formation processes are under ongoing investigation. OA can be directly emitted as primary particulate matter (primary OA, POA) or formed through the oxidation of gas-phase precursors with subsequent condensation or nucleation (secondary OA, SOA). SOA dominates submicron OA at remote sites (90%) and is also a substantial contributor (30-80%) in the urban environment (Zhang et al., 2011).



Mass spectrometry has significantly advanced the chemical characterization and quantification of OA. Instruments equipped with electron-impact (EI) ionization, (EI), such as the aerosol mass spectrometer (AMS, Canagaratna et al., 2007) and aerosol chemical speciation monitor (ACSM, both Aerodyne Research, Inc., Ng et al., 2011a; Fröhlich et al., 2013) provide quantitative online measurements of OA (Jimenez et al., 2016). The application of positive matrix factorization (PMF) to AMS and ACSM mass spectra has allowed the separation of different POA sources such as traffic, cooking, and wood burning, as well as oxygenated OA factors representing SOA (e.g. Jimenez et al., 2009; Lanz et al., 2007; Lanz et al., 2008; Lanz et al., 2010; Crippa et al., 2014; Canonaco et al., 2013). While different SOA factors identified by AMS-PMF have been separated according to their degrees of oxygenation and volatility, information on the different origins of this fraction is limited due to significant fragmentation of the molecules by EI. Moreover, SOA from different sources converges to a chemically similar composition during oxidation (Kroll et al., 2011, Ng et al., 2011b).

Other strategies for aerosol mass spectrometry have provided complementary information. Ionization by electrospray (ESI) avoids significant analyte fragmentation. The coupling of ESI to ultra-high-resolution Fourier-transform mass spectrometers thus provides detailed information on the chemical composition of a sample. However, using ESI not all compound classes can be detected efficiently due to ion suppression (Furey et al., 2013; Trufelli et al., 2010; Kourtchev et al., 2013). In addition, high costs and labor intensity restrict the application typically to smaller sets of samples. Traditional techniques such as gas- or liquid-chromatography (Hoffmann et al., 2011) coupled to MS suffer from similar restrictions.

Laser-based mass spectrometers such as laser-desorption/ionization mass spectrometers (LDI-MS) and two-step laser mass spectrometers (L2MS) are, similarly as ESI, less affected by fragmentation than EI. Haeffliger et al. (2000a and 2000b) show that such instrumentation allows for identification of the dominant primary sources using mass spectral fingerprints (tracer  $m/z$ 's identified using emission samples and principle component analysis of ambient measurements). LDI-MS instruments may be subject to matrix effects, i.e. ion formation from a given compound does not only depend on its abundance but also on the abundance of all other compounds (De Hoffmann and Stroobant, 1999). Additional matrix sensitivities may arise if ion-neutral reactions take place in the desorbed plume, which means that MS design (e.g. time-of-flight vs. ion trap) also plays a role (Murphy, 2007). Although the importance of matrix effects is evident, studies with a single-particle LDI-MS (ATOFMS, laser wavelength 266 nm) have achieved good correlations of major aerosol components with well-established reference measurements like such as elemental carbon (EC, Sunset Analyzer, Sunset Laboratory Inc.), OA (AMS), NH<sub>4</sub> (AMS), SO<sub>4</sub> (AMS), NO<sub>3</sub> (AMS), and K (collected with a particle-into-liquid sampler, PILS), suggesting that underlying matrix effects did not dominate measurement reproducibility (Healy et al., 2013).

Online Similar to the AMS, online single-particle LDI-MS instruments, like such as the ATOFMS, generally employ very high laser fluences. This leads to similar also yield extensive fragmentation as observed in the AMS. However, such fragmentation can be avoided. By measuring offline aerosol samples (filters) at different laser powers (LDI-MS, Shimadzu/Kratos, Axima CFR) and with and without adding a matrix, using other systems. Samburova et al. (2005a) showed by comparing measurements with and without matrix addition, that fragmentation was negligible in their instrument. (wavelength 337 nm, LDI-MS, Shimadzu/Kratos, Axima CFR). Overall, offline LDI-MS may therefore provide quick access to additional chemical information at near-to-molecular level, potentially allowing differentiation between several primary organic aerosol sources and even different precursor-related SOA categories (Kalberer et al., 2004; Samburova et al., 2005a). Based on offline LDI-MS measurements performed for an urban background site in Zurich, Switzerland (same site as Haeffliger et al., 2000a; 2000b), Baltensperger et al. (2005) suggested that SOA from biogenic precursors is more important than SOA from anthropogenic precursors. However, such LDI-MS analyses are rare, typically focusing on pattern analysis in the mass spectrum since LDI-MS signal quantification is difficult due to the variability in ionization efficiencies and

chemistry for different compounds. In addition, LDI-MS has not been applied to extensive datasets. In comparison to online analyses, analyzing offline filter samples allows covering longer time periods and larger observation networks.

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In this work, we evaluated the use of offline LDI-MS for the direct measurements of PM collected on filters, without the addition of an ionization matrix. We assessed the contributions of different aerosol sources to the total PM. To control for instrumental differences, we measured all filters on the same instrument. We developed novel procedures for instrument calibration and uncertainty assessment. We measured three source reference samples (traffic, wood burning, and cooking) in addition to 819 ambient samples. The ambient samples included filters from the entire year of 2013 at nine sites in central Europe with different emission conditions (Alpine valleys strongly influenced by wood burning, as well as urban and rural regions). Based on positive matrix factorization (PMF) using LDI-MS mass spectral data, the ability to resolve OA sources in source apportionment was assessed and used for obtaining a deeper understanding of sources contributing to the organic aerosol in central Europe.

## 2. Methods

### 2.1 Sample collection and other chemical analyses

Ambient samples were collected at nine sites in Switzerland and Liechtenstein (described in more detail in Daellenbach et al., 2017) covering different atmospheric conditions (urban/rural and background/curbside). Seven sites were located on a plateau north of the Alpine crest (Basel, Bern, Payerne, Frauenfeld, St. Gallen, Zurich, Vaduz); the remaining two sites were located in Alpine valleys south of the Alps (Magadino, San Vittore). Samples were collected on quartz filters (Pall Corp.) by local air quality monitoring networks every 4<sup>th</sup> day during the entire year 2013 (819 filters) and used for gravimetric PM<sub>10</sub> quantification. The samples were stored at -18°C and transported in cooling boxes. Before any further handling steps, the samples were allowed to warm up at room temperature for 60 minutes in order to avoid condensation of ambient humidity. The organic (OC) and elemental (EC) carbon content was determined by a thermo-optical transmission method using a Sunset OC/EC analyzer (Sunset Laboratory Inc., Birch and Cary, 1996), following the EUSAAR-2 thermal-optical transmission protocol (Cavalli et al., 2010). For 33 samples from Magadino, radiocarbon (<sup>14</sup>C) analyses were performed (2013; 10 samples, 2014; 23 samples, Vlachou et al., in prep). The radiocarbon analyses were conducted at the University of Bern at the Laboratory for the Analysis of Radiocarbon with the Accelerator MS (LARA; Szidat et al., 2014) to determine the contribution of fossil and non-fossil OC (OC<sub>f</sub> and OC<sub>nf</sub>, respectively) and EC (EC<sub>f</sub> and EC<sub>nf</sub>, respectively) content (Zhang et al., 2012). Major ion concentrations were measured by ion chromatography (IC) by a Dionex ICS1000 instrument (Piazzalunga et al., 2013; Jaffrezo et al., 1998). Levoglucosan was measured by high-performance anion exchange chromatography (HPAEC) with pulsed amperometric detection (PAD) using an ion chromatograph (Dionex ICS1000) following the method of Piazzalunga et al. (2010; 2013a). Gas-phase NO<sub>x</sub> was measured online using a chemiluminescence method, and meteorological parameters (e.g., temperature) were monitored. Further, equivalent black carbon (eBC) was measured with a multi-wavelength Aethalometer AE 31 (Magee Scientific Inc.) (Hansen et al., 1984; Herich et al., 2011) in Magadino, Payerne and Zurich. eBC was separated into a wood burning influenced (eBC<sub>wb</sub>) and traffic influenced (eBC<sub>tr</sub>) fraction based on the enhanced absorption of eBC<sub>wb</sub> in the ultraviolet range. For this computation, we used the Ångström exponents for wood-burning ( $\alpha_{wb}$ ) of 1.7 and for traffic ( $\alpha_{tr}$ ) 0.9 (Zotter et al., 2017; Sandradewi et al., 2008). On all the samples, offline AMS analyses (Daellenbach et al., 2016) were conducted. This involved the analysis of the water-soluble organic matter (WSOM) by a high resolution time-of-flight AMS. The offline AMS data were used for source apportionment (Daellenbach et al., 2017; Bozzetti et al., 2016) and used in this study for comparison; the source mass concentrations were corrected for the missing water-insoluble mass fraction using the method described in Daellenbach et al. (2016).

Reference source samples for traffic (representing PM<sub>10</sub>) were collected in a tunnel (Islisberg tunnel on Swiss highway A4, exit Wettswil am Albis). The PM<sub>10</sub> samples were collected on Sunday, 2014-05-18, and Tuesday, 2014-05-20, from midnight to noon with 2 hour intervals (rush hour in the morning). Additional source samples were collected on quartz filters for beech-wood burning through laboratory experiments of whole cycle burns and stable flaming phase only. These samples did not only include the primary aerosol, but also two different levels of aging, i.e., after simulated aging in a smog chamber for 1 hour (equivalent to an OH dose of 10<sup>7</sup> cm<sup>-3</sup> h ) and 4 hours (OH dose 3·10<sup>7</sup> cm<sup>-3</sup> h) (Bruns et al., 2015; 2016). Additionally, also filter samples collected during cooking experiments were analyzed (Klein et al., 2016a, 2016b).

## 2.2 Laser-desorption/ionization ToF MS analysis

### 2.2.1 Instrument and measurement settings

In laser desorption/ionization mass spectrometry, LDI-MS, liquid or solid material is simultaneously desorbed and ionized by pulsed laser irradiation. The laser is focused on the surface of the sample. Ions are produced through two major pathways (Knochenmuss et al., 2000; Zenobi et al., 1998). In the first pathway, ions are formed by interaction with the laser beam (primary ions). In the second pathway, in the expanding primary plume of desorbed molecules and primary ions, ion-molecule reactions produce secondary ions (Knochenmuss et al., 2000). The detailed ionization mechanisms are still under investigation (Knochenmuss et al., 2002; 2003; 2006; 2016). Although this type of ionization is considered softer than electron ionization, the observed ions are typically still fragments of their parent molecules (De Hoffmann and Stroobant, 1999). Furthermore, the measured intensity of one compound does not only depend on its concentration but also on the concentration of all other compounds present (the so-called matrix effect) making quantification challenging (Ellis et al., 2014; Borisov et al., 2013).

We recorded the mass spectra of 819 filter samples (in an  $m/z$  range 65-500,  $\text{thomsons (Th; 1 Th = 1 Da } e^{-1}$ , where  $e$  is the elementary charge, ion gate at  $m/z$  60 Th) using a laser-desorption/ionization-ToF MS (Shimadzu Axima Confidence, Shimadzu-Biotech Corp., Kyoto, Japan) equipped with an N<sub>2</sub> laser (wavelength 337 nm, frequency 50 Hz, laser pulse width 3 ns, 130-180  $\mu\text{J/pulse}$ ) in the positive reflectron mode. All the accessible instrumental parameters were kept constant during the whole period of measurements taking place from November 2015 to mid-March 2016. Specifically, the laser intensity was adjustable by means of a rotating wheel of filters with varying transmission. While this wheel was initially set transmissivity (0 being blocked and 180 being completely open). We set the wheel parameter to 105 of 180 which would result in an estimated laser energy of  $\sim 6\text{-}9 \mu\text{J/pulse}$  or  $2.8\text{-}4.2 \cdot 10^8 \text{ W/cm}^2$  with a 3ns pulse and 30 $\mu\text{m}$  laser beam diameter. While this wheel the laser energy was initially set (wheel parameter 105 of 180 equaling roughly 6-9  $\mu\text{J/pulse}$  or  $2.8\text{-}4.2 \cdot 10^8 \text{ W/cm}^2$  with a 3ns pulse and 30 $\mu\text{m}$  laser beam diameter, 0 being blocked 180 being completely open) and kept constant, the aging of the laser during the given time period was also expected to reduce its intensity. We monitored and assessed changes in laser power and other instrumental parameters, as well as possible sources of uncertainty/contamination from sample preparation and intra and inter-day reproducibility, by repeated measurements of a subset of our samples.

Quartz filter punches of 8 mm were attached to a custom-made stainless steel sample holder (32 slots). Each of the samples was additionally spiked with a droplet of dilute aqueous silver nitrate solution (AgNO<sub>3</sub>, Sigma Aldrich, >99.8%, 500 ppt to 20 ppm), to provide Ag cations as an internal standard for the  $m/z$  calibration. After drying under ambient conditions, filter punches were inserted into the sampling chamber and analyzed by the LDI-MS. Blanks were measured according to the same procedure. Both intra-day and delayed repeated measurements were conducted for the same filters to assess instrument performance and uncertainty/contamination from sample preparation. The intra-day repeatability was assessed by measuring

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3 filters 10 times on 3 different occasions as outlined in Section 3.1. Overall, 96% of the available ambient filter samples (785 filters) provided usable data (defined below).

## 2.2.2 Data treatment

While most atmospheric LDI-MS studies present raw mass spectra (Samburova et al., 2005a; 2005b; Kalberer et al., 2004; Baltensperger et al., 2005), in the present study, we introduced data treatment techniques in order to perform further mass spectral analysis on stick-integrated spectra. The techniques we employed are described in the following.

The  $m/z$  calibration ~~was performed using silver nitrate solution~~ parameters were highly variable between different samples. Therefore, to determine unit mass resolution integration regions it was necessary to perform an  $m/z$  calibration on every single sample. However, unlike in the Aerodyne AMS ( $\text{N}_2^+$ ,  $\text{O}_2^+$ ,  $\text{W}^+$ ), there are no dominant anchor ions present in these spectra that could be used for an  $m/z$  calibration. In absence of such ions, we performed a two-step calibration procedure as described in the following. Each sample was spiked with silver nitrate ( $\text{AgNO}_3$ , Sigma Aldrich, >99.8%) as an internal standard (approach illustrated in Fig. 1). In order to avoid the suppression of the sample signal, the internal standard (aqueous solution 500 ppt to 20 ppm) was only placed (as a droplet) on a small part of the sample. The 499 spectra from all positions on the measurement grid were separated and defined as (1) silver-spiked (lower panel in Fig. 1a), (2) silver-free (upper panel in Fig. 1a), and (3) intermediate-silver. Intermediate-silver cases were defined using the signal intensity in the regions of the mass spectrum where silver was expected, in comparison to adjacent silver-free regions of the spectrum, and were discarded.

~~In an initial~~ To calculate the first calibration, we calibrated the average silver-spiked spectrum of each filter sample, using the peaks of the silver monomer ( $m/z$  107, 109 ~~Th~~), dimer ( $m/z$  214, 216, 218 ~~Th~~), and trimer ( $m/z$  321, 323, 325, 327 ~~Th~~). We found that this calibration of silver-spiked was not directly applicable to the average silver-free mass spectrum. Possibly, spiking the filter region with aqueous  $\text{AgNO}_3$  caused enough of a change to the surface of the filter sample to influence the ionization physics. This could affect the  $m/z$  calibration, since ~~the delayed-pulse~~ ion extraction in our instrument is not orthogonal to the ablation plume, ~~but nearly parallel~~. Therefore, a second calibration was obtained for the averaged silver-free spectra ~~were recalibrated with respect to the silver-spiked spectra~~ using prominent non-silver peaks present in both spectra. Such a two-step calibration is necessary to achieve accurate  $m/z$  calibrations of the silver-free spectra.

After calibration, mass spectra were baselined using the following custom algorithm. A window of width 1 u (u being the unified atomic mass unit, the detector bins were approximately 0.04 u in width) was applied to identify the lowest signal intensity of that range, which was defined as the baseline. The window was moved across the mass spectrum with steps of 0.25 u to obtain a baseline spectrum. After linear interpolation, the baseline was subtracted from the spectrum. Subsequently the spectra were integrated to unit mass resolution sticks (UMR, 1 stick per unit mass). The UMR integration window was not centered at integer masses, but rather at integer mass plus the mass defect of an alkyl ion ( $\text{R-CH}_2^+$ ). The width of this window was defined as extending to the minimum signal intensity above and below the center using the first derivative of the signal.

To enable a direct comparison of LDI-MS and HR-AMS results, the intensity ~~scales~~ of the calibrated and baseline-subtracted spectra ( $I'_{m/z}$ ) with arbitrary units were scaled to OM ( $\mu\text{g}/\text{m}^3$ ) using the respective OM contents of the filter samples (see Eq. 1, with rescaled intensities termed  $I_{m/z}$  thereafter).  $\text{OC}_{\text{Sunset}}$  was determined using the Sunset OC/EC analyzer and  $(\text{WSOM/WSOC})_{\text{oAMS}}$  through WSOM measurements (described in Section 2.1), ~~WSOC being water-soluble organic~~

[carbon](#)). Note that an additional contribution of non-organic material to the signal could not be fully excluded. However, we consider the detected signal mainly to be related to organics based on the measurements of ambient aerosol with offline LDI-MS systems using different laser-wavelengths (193 nm and 355 nm, Aubriet et al., 2010) and pure components (337 nm, Goheen et al., 1997) indicating that sulfate and nitrate respond only in the negative mode. Additionally, molecular nitrate and ammonium ions are too small to be detected in this study (cutoff at  $m/z$  65 [Th](#)). We also do not expect clusters between inorganic cations and organics because we did not observe significant amounts of silver/organic clusters in silver-spiked spectra (Fig. 1a).

$$I_{m/z} = \frac{I'_{m/z}}{\sum_{m/z} I'_{m/z}} * OC_{\text{sunset}} * \left( \frac{WSOM}{WSOC} \right)_{oAMS} \quad (1)$$

## 2.3 Source apportionment / PMF

### 2.3.1 General and input data

Positive matrix factorization (PMF, Paatero and Tapper, 1994; Paatero, 1997) is a widely used algorithm for source apportionment. PMF (Eq. 2) explains the variability in a dataset ( $x_{i,j}$ , here LDI-MS mass spectra scaled to OM,  $I_{m/z}$ ) as a linear combination of constant factor profiles ( $f_{k,j}$ , here mass spectral signatures) and their time-dependent contributions ( $g_{i,k}$ , here concentrations of this factor). The index  $i$  represents a specific point in time,  $j$  the signal at a specific  $m/z$ , and  $k$  a factor (up to the number of factors  $p$ ).

$$x_{i,j} = \sum_{k=1}^p g_{i,k} f_{k,j} + e_{i,j} \quad (2)$$

PMF is solved by minimizing the residuals  $e_{i,j}$  weighed by the measurement uncertainty, here  $\sigma_{PMF,i,j}$ . In this study, PMF was solved by the multilinear-engine 2 (Paatero, 1999 and references therein) using the front-end user interface SoFi 4.9 (Canonaco et al., 2013) developed for Igor Pro v6 (Wavemetrics). We performed PMF without constraints of e.g. reference mass spectra of aerosol sources (unconstrained PMF). The data matrix consisted of the stick-integrated LDI-MS spectra ( $m/z$  65 to 485 [Th](#) scaled to OM) of all 785 filter samples. We excluded all  $m/z$ 's for which we expected silver signals (listed in section 2.2.2) even though only silver-free spectra (according to the definition given above) were considered.

We determined the PMF uncertainty matrix ( $\sigma_{PMF,i,j}$ ) based on the instrument repeatability (described in Section 3.1) according to the two-component model for measurement uncertainty described by Rocke and Lorenzato (1995) and Wilson et al. (2004). This approach ultimately means that an absolute-error term (e.g., due to noise) is combined in quadrature with a relative-error term (e.g., due to scaling or calibration uncertainties) as shown Fig. 2 and in Eq. 3 ( $\sigma'_{PMF,m/z}$ , for a single mass spectrum). The need for and applicability of such an error model has been demonstrated by Corbin et al. (2015) for mass-spectrometry-PMF. We obtained the two error terms by fitting Eq. 3 to the relative standard deviation of replicate measurements for each filter and experiment, and applied the averaged fitted parameters to the entire data set (constant error term,  $\sigma_{\text{abs}}$ , and an error term proportional to the measured peak intensity,  $I'_{m/z} \sigma_{\text{rel}}$ ).

$$\sigma'_{PMF,m/z} = \frac{\sqrt{\sigma_{\text{abs}}^2 + I'^2_{m/z} \sigma_{\text{rel}}^2}}{I'_{m/z}} \quad (3)$$

The error was scaled to OM in the same way as the LDI-MS mass spectra (Eq. 4).

$$\sigma_{\text{PMF},m/z} = \sigma'_{\text{PMF},m/z} \left( \frac{I_{m/z}}{I'_{m/z}} \right) \quad (4)$$

### 2.3.2 Uncertainty estimate of PMF results

A well-established statistical tool for estimating the uncertainty is the bootstrap technique, which consists of randomly resampling the input data, with replacement, to create input matrices with the same dimensions as the initial data matrix (Davison and Hinkley, 1997; Brown et al., 2015). We performed 1'000 bootstrap PMF runs (also using the replicate measurements) with different initial guesses ("seeds") using input matrices of the same size (785 filter samples and 412 ions), resulting in the uncertainty estimate  $\sigma_{\text{bs}}$ . We also assessed and parameterized the uncertainty arising from the intra-day repeatability for three filters analyzed repeatedly by LDI-MS on 3 different days, 10 times each day ( $\sigma_{\text{intraday}}$ , see SI). The reported uncertainty for each PMF factor  $\sigma_{\text{tot}}$  is the quadratic sum of  $\sigma_{\text{bs}}$  and  $\sigma_{\text{intraday}}$ , as explained in more detail in the Supplement. Potential long-term drifts in instrumental response were evaluated ( $\sigma_{\text{interday}}$ , see Supplement, Fig. S10) but were not accounted for numerically.

## 3. Results and Discussion

### 3.1 Calibration, repeatability and quantification

An example of a sample spectrum after  $m/z$  calibration and baseline subtraction, separated into silver-spiked and silver-free spectra and averaged over the entire filter punch, is presented in Fig. 1a. The silver-spiked spectrum features an intensive signal at  $m/z = 214.214$  Th related to the  $^{107}\text{Ag}_2$  dimer, in contrast to the silver-free spectrum.

The operational mass resolution,  $m/\Delta m$ , based on the measurements conducted on one sample tray per month and the accuracy of both calibration steps are shown in Fig. 1b and 1c, respectively. The operational resolution of the instrument was only determined for the silver mono- (resolution 1100), di- (1700), and trimer (2100) as we are confident of the absence of interfering ions for these peaks. In comparison to the LDI-MS, the HR-ToF AMS in V-mode (and W-mode) has a resolution of ca. 2'000 (4'000) at  $m/z$  100 Th and an  $m/z$  calibration accuracy  $< 20$  ppm ( $< 10$  ppm, DeCarlo et al., 2006). With this accuracy and resolution, neither distinguishing different ions at the same nominal mass ~~was not~~ nor estimating properties such as O/C and H/C for a nominal mass following Stark et al. (2016) was possible for the quartz-filter-LDI-MS measurements. Instead, the spectra were integrated to UMR sticks. The presence of considerable signal at odd masses might indicate that significant fragmentation occurs during desorption and ionization of the organic matter in the LDI-MS (Fig. 1a, 3, 4). However, Samburova et al. (2005a) suggested that fragmentation in this instrument is negligible. The mass spectral signatures acquired in this study are similar to the ones in Samburova et al. (2005a) and the lower laser energy applied in this study is lower than in the former study. This suggests that also during our measurements fragmentation was not a prominent process. Overall, the extent of fragmentation remains unclear.

The repeatability of the mass spectral signature was assessed in detail for 3 selected filters (collected in Basel on 2013-06-21, 2013-09-21, and 2013-12-22). Ten punches from each of these 3 samples were prepared on the sample holder in the same time and subsequently measured. The given procedure was repeated on 3 occasions (on 2015-11-25, 2015-12-17, 2016-01-28). The relative error was related to the absolute signal intensity for each  $m/z$  for each sample and experiment (Fig. 2), with an asymptotic value of only  $\sim 9 \pm 7$  % at high signals (average and standard deviation of fitting Eq. 3). The average absolute error for small signals was  $102 \pm 48$  a.u. (from Eq. 3). These parameters did not show a temporal trend which suggests that the measurements are repeatable, despite filter inhomogeneity and laser instability. Analyses of field blanks (also spiked with

AgNO<sub>3</sub>) exhibited low signal with 99% of peaks below detection limit (defined as  $3 * \sigma_{\text{abs}}$ ). The other 1% of the peaks ( $m/z$  197, 249, 251, 322, 324 [Th](#)) exhibited high variability among the field blank analyses (see details in Supplement, Fig. S1 and S2). Therefore, no blank subtraction was performed.

The total measured intensity ( $I_{\text{tot}} = \sum I_{m/z}$ ) did not show a relation with the filter loading (OC, OC+EC, PM<sub>10</sub>, see Supplement, Fig. S3), indicating that factors such as the composition, size, and/or mixing state of the collected aerosol particles had a larger influence on the measured ion intensities than the mass of PM. Moreover, replicate measurements indicated that the  $I_{\text{tot}}$  of a single sample measured on different instances was affected by an instrumental drift (details in Supplement, Fig. S4).

### 3.2 Combustion-source samples

Filters from cooking processes (frying, related largely to oil pyrolysis, Klein et al. 2016a; Allan et al., 2010) did not yield measurable mass spectra in our LDI-MS. This observation is in agreement with a prior study which performed in-situ single-particle LDI-MS measurements at 266 nm (using the ATOFMS) and did not observe cooking particles (Healy et al., 2013).

The lack of either graphene-like black carbon or polycyclic aromatic hydrocarbons (PAHs) to absorb the LDI-MS laser may explain these observations (Healy et al., 2013), which suggests that our measurements are more sensitive towards certain PM components. We speculate that it may be possible in some cases to observe cooking-related ions in LDI-MS of atmospheric samples, if such ions are secondary ions formed in the ablation plume or cooking particles are coated by absorbing SOA.

The mass spectra of tunnel filter samples were complex and characterized by distinct patterns (Fig. 3). Some  $m/z$  fragments (e.g. 84, 94, 101, 177 [Th](#)) were present in the weekday and weekend tunnel samples, but not in the biomass smoke. Between the two tunnel samples, large differences could also be identified. The tunnel sample collected on a Tuesday during rush hour (06:00-08:00) showed many more peaks (among others 163 and 177) which were less prominent during the weekend at the same time and in the same location (where  $m/z$  101 and 143 [Th](#) dominated the weekend spectrum).

Previous studies have identified lubrication oil as a major component of diesel- and gasoline-vehicle exhaust (Gentner et al., 2017; Chirico, et al., 2011), which consists largely of aliphatic hydrocarbons. Cooking particles similarly consist of such hydrocarbons (Schauer et al., 2002; Allan et al., 2010; Crippa et al., 2013). Since we did not observe a mass spectrum for the cooking sample, but do observe one for the tunnel sample, our detection of tunnel particles may rely either on the presence of black carbon or PAHs on the filter samples. To the extent that these species are heterogeneously distributed within the sampled aerosol particles, our technique may be specifically sensitive to certain components of traffic emissions.

The mass spectrum of the wood burning sample from the whole burning cycle shows a complex pattern with a bimodal envelope of relatively-intense ions from  $m/z$  200 [Th](#) onwards. By contrast, the sample from the stable-flaming phase shows fewer peaks (e.g.  $m/z$  85, 140, 213 [Th](#), etc.) with a less significant envelope of ions.

### 3.3 Ambient samples

Typical mass spectra from 4 different locations, known to be heavily influenced by specific sources at different times of the year are displayed in Figure 4. The fragments are color-coded by their correlation coefficients with 3 different markers or proxies for aerosol types, i.e., NO<sub>x</sub> for traffic exhaust (Fig. 4a), levoglucosan for biomass burning emissions (Fig. 4c and 4d) and temperature for SOA production (as temperature is expected to exponentially enhance biogenic emissions, Leaitch et al., 2011, Fig. 4b). In general, the mass spectra from winter-time were characterized by higher contributions from higher-



molecular-weight fragments compared to summer. For the winter sample from San Vittore, a large “hump” was present, on top of which signals appeared with high intensities and a highly regular pattern of  $m/z$  differences of 14 Th. These fragments, albeit not all detected in smog chamber biomass burning aerosols, were highly correlated with levoglucosan, a unique marker of biomass smoke.

The  $m/z$  84 94, 101, 120, 143, 165, and 177 Th, also detected in tunnel samples, showed high relative intensities in Bern in summer, of which 84, 94, 120, and 177 Th correlated with  $\text{NO}_x$ , which suggests a high contribution of traffic emissions at this location, consistent with previous observations (Zotter et al., 2014). These ions were also observed, but to a lesser extent, at the rural site of Payerne. At the latter site, the highest intensities were found for  $m/z$  74 and 104 Th. Only these 2 ions showed a clear relation with temperature and a clear increase during summer. Overall, this spatial and temporal variability observed in the spectral fingerprints and the correlation of specific fragments with environmental parameters suggest that LDI-MS data contain mass spectral information that may be used to separate the contribution from different aerosol sources. Source apportionment analysis using these data in combination with PMF was therefore explored in the following.

### 3.4 Source apportionment results

#### 3.4.1 PMF setup

Residual analysis of preliminary PMF runs showed that structure in the residuals was removed when increasing the number of factors up to 5 but not when further increased. However, when allowing for 7 factors, further environmentally interpretable separations could be achieved. Increasing to 8 factors led to the separation of a third traffic related factor, which however contributed little to the overall mass. Therefore, we opted for 7 factors (see details in Supplement, Fig. S6, S7, and S8).

#### 3.4.2 Interpretation of PMF factors

##### 3.4.2.1 Traffic-related factors

Two factors could be related to traffic (Fig. 5a and 5b, termed traffic1 and traffic2, respectively) based on patterns in the factor profiles similar to patterns in the samples measured in the tunnel on a weekday (Fig. 3a). Since the signatures obtained from the tunnel filters represent both tailpipe exhaust and resuspended dust, the two factors identified could be a mixture of both sources. In order to elucidate the reasons behind the separation of traffic related aerosols into two factors by PMF, we inspected their relationship with  $\text{NO}_x$  and  $\text{eBC}_{\text{tr}}$ , typical markers of traffic emissions. Both traffic1 and traffic2 showed a ~~relationship to increasing concentrations with increasing  $\text{NO}_x$ , only levels.~~ However traffic1/ $\text{NO}_x$  and traffic2/ $\text{NO}_x$  were seasonally variable ( $R_{p,\text{traffic1},\text{NO}_x}=-0.17$ ,  $R_{p,\text{traffic2},\text{NO}_x}=0.24$ , Fig. 6, Fig. S9). While traffic2 correlated with  $\text{eBC}_{\text{tr}}$ , traffic1 did not show such a dependency (Tab. S1). ~~This, Fig. 6, Fig. S9).~~ The correlation with  $\text{eBC}_{\text{tr}}$  suggests that traffic2 is related to primary emissions from the combustion process ~~and that traffic2 is.~~ Based on the lack of correlation between traffic1 and  $\text{eBC}_{\text{tr}}$ , traffic1 might also be influenced by other processes, likesuch as e.g. secondary production. Additionally, tailpipe exhaust cannot be distinguished from other traffic related emissions in the tunnel samples and might also contribute to traffic1.

The ratio of both factors to  $\text{NO}_x$  exhibited a clear seasonality, increasing during summer (Fig. 6a and 6b). Such a change can either be caused by a change in the emission patterns, by an enhanced photochemical production of these factors or due to a



change in the lifetime of e.g.  $\text{NO}_x$ . In Fig. 6c and d, we compared the contribution of the traffic related factors with those of  $\text{eBC}_{\text{tr}}$ , another tracer of traffic emissions whose lifetime is similar to OA (see Fig. S9 for summer points only). The comparison shows that while the ratio of traffic2 to  $\text{eBC}_{\text{tr}}$  is not season-dependent, the one of traffic1 to  $\text{eBC}_{\text{tr}}$  increases during summer. The seasonal variability in the relative contribution of traffic1 might thus be related to a seasonal change in fleet composition or combustion conditions, or an enhancement of the photochemical production of this fraction. An enhanced dust resuspension in the warm season (e.g., more dust on the road due to less precipitation) could also contribute to an increased traffic1 concentration in summer.

### 3.4.2.2 Wood-burning-related factors

Three factors could be related to wood burning emissions (Fig. 5c, 5d, 5e). The first (Fig. 5c) showed a similar mass spectral pattern as samples from laboratory experiments from stable flaming wood burning exhaust from a log wood burner and thus was termed efficient wood burning (BBeff). The fingerprint of the second (Fig. 5d) resembled that of laboratory wood smoke aerosols from the entire burning cycle including the inefficient starting and burnout phase and, therefore, was related to inefficient wood burning (BBineff1). BBineff1 showed also a similar signature as spectra from wood burning haze episodes in San Vittore identified by high levoglucosan concentrations. The third (Fig. 5e) mostly explained the masses above  $m/z$  300 [Th](#). The signature of the whole cycle wood burning aerosols from a log wood burner (Fig. 3c) also showed high relative contributions at high  $m/z$ , similar to this factor. For this reason this factor was termed BBineff2. Both BBeff and BBineff1 correlated with levoglucosan (Fig. 6e, 6f, Tab. S1). Similar to BBeff and BBineff1, BBineff2 correlated with levoglucosan and thus could be related to wood burning emissions (Fig. 6g, Tab. S1). However, BBineff2 also correlated with  $\text{NH}_4^+$ , a marker of aged aerosols, at the northern sites (Fig. 6m), which may also indicate that this fraction can include aged and/or secondary components (Tab. S1).

Comparing the wood burning related factor time series to potassium ( $\text{K}^+$ ), an inorganic wood burning marker mostly present in ash, provides further insight into the separation of the 3 wood burning related factors. Among the factors related to wood burning emissions, BBeff has a lower  $\text{BB}/\text{K}^+$  ratio (2.4, IQR 1.25-4.2) than BBineff1 (6.0, IQR 2.7-12.0) and BBineff2 (11.6, IQR 6.0-21.0). In wood burning experiments, it was found that OC, EC,  $\text{PM}_{10}$ , and PAH emissions increase relative to the potassium output during non-ideal burning conditions (Lamberg et al., 2011). Zotter et al. (2014) found a north-south gradient in levoglucosan/ $\text{K}^+$  and  $\text{OC}_{\text{nf}}/\text{K}^+$  in Switzerland and hypothesized that it might be linked to the burning conditions. Thus the higher the LDI-BB/ $\text{K}^+$  ratios (north: 16.6, IQR 10.4-30.8 south: 30.8, IQR 22.5-44.2), the less efficient the burning conditions which in turn supports the hypothesis that BBeff represents the most efficient burning conditions among the three factors. All wood burning related factors showed higher median  $\text{BB}/\text{K}^+$  ratios at the southern Alpine valley site (BBeff: 3.0 with IQR 1.1-5.2, BBineff1: 10.9 with IQR 7.5-17.3, BBineff2: 15.8 with IQR 9.9-24.0) than in northern Switzerland (BBeff: 2.1 with IQR 1.2-3.4, BBineff1: 4.3 with IQR 2.2-8.6, BBineff2: 9.6 with IQR 5.2-18.3) as also visible in Fig. 6h/i/j. The north-south gradient in the  $\text{BB}/\text{K}^+$  ratios might be caused by imperfections of the separation of the burning conditions by PMF, but other effects [like such as](#) the age of the stove population and the used technology could contribute as well. The wood consumption (BFE, 2013) of automatic burners (>50kW), was higher in northern Switzerland (in the respective regions,  $48\text{m}^3$  wood/ $\text{km}^2$  area) than in southern Switzerland ( $8\text{m}^3$  wood/ $\text{km}^2$  area). Since the OM, POA, and SOA emissions of pellet burners during the stable phase are drastically reduced compared to modern logwood burners in a stable flaming phase (e.g. Heringa et al., 2011), they might contribute over-proportionally to potassium in northern Switzerland but only little to OM, leading to the higher  $\text{BB}/\text{K}^+$  ratios at the southern Alpine sites.

### 3.4.2.3 Biogenic-~~SOA~~OA, low-molecular-weight OA, and other OA factors

A factor characterized by high contributions of low-molecular weight ions (Fig. 5f, LMW-OA) correlates with  $\text{NH}_4^+$  suggesting secondary processes as origin (Fig. 6l, Tab. S1). At the southern Alpine valley sites, the LMW-OA /  $\text{NH}_4^+$  ratio was higher than in northern Switzerland. As  $\text{NH}_4^+$  is mostly associated with the secondary inorganic species sulfate and nitrate this variability will mostly be related to differences in the VOC versus  $\text{SO}_2$  and  $\text{NO}_x$  emissions, along with temperature differences influencing the partitioning of nitrate to the particle phase.

The contribution of the last remaining factor showed an exponential increase with temperature, similar to terpene emissions and biogenic SOA (Leaith et al., 2011), suggesting this factor to be strongly influenced by biogenic SOA production.

Therefore, this factor was termed bio-OA. The relationship between bio-OA and temperature was similar both at the northern and southern sites (Fig. 6k, Tab. S1). Bio-OA had a highest relative contribution at the most rural site in the dataset (Payerne: yearly average bio-OA 11%, yearly average  $\text{NO}_x$  concentration: 9 ppb) and the lowest at the most trafficated site (Zurich: yearly average bio-OA 3%, yearly average  $\text{NO}_x$  concentration: 48 ppb). The chemical signature of bio-OA was dominated by fragments at  $m/z$  74 and 104 Th (Fig. 5g). The nature of these fragments remained unidentified, but could not be related to ~~mono-terpene~~monoterpene or sesquiterpene SOA because of their very low  $m/z$ . Further, the composition of secondary aerosols is expected to be much more complex, showing a series of fragments distributed over a wide  $m/z$  range.

We note that the abundance of the fragments at  $m/z$  74 and 104 Th relative to the total signal should not be directly related to their absolute concentrations and their ionization efficiency might be superior to that of other molecules, which may increase their apparent contribution.

### 3.4.3 Comparison to offline AMS ~~results and earlier AMS campaigns~~and assessment of LDI-MS response factors

The LDI-MS source apportionment results were compared to those based on offline AMS (oAMS) analyses performed on the same samples (Daellenbach et al., 2017). We note that this comparison is not straightforward as different sources are separated by the two methods. For this purpose traffic1 and traffic2 were summed up to LDI-traffic, and BB<sub>eff</sub>, BB<sub>ineff</sub>1, and BB<sub>ineff</sub>2 to LDI-BB (Fig. 7).

As was the case for the comparison to  $\text{NO}_x$  and  $\text{eBC}_{\text{tr}}$  (Section 3.4.2), also in the comparison to the offline AMS traffic (HOA), a higher LDI-traffic/HOA was observed in summer which contributes to the low correlation coefficient ( $R^2=0.0504$ ). This might be related to the fact that HOA is only primary without major contributions of traffic SOA or dust resuspension while LDI-traffic is potentially also influenced by aged/secondary traffic aerosol and resuspension. In contrast to summer, we observe a good agreement between LDI-traffic and HOA in winter. Thus even though a varying relative response factor (rRF) of the LDI-MS might contribute to these differences, these biases are not systematic but season dependent. As stated earlier, LDI-traffic is thought to be a mixture of primary tailpipe exhaust, aged/secondary tailpipe exhaust, and resuspended dust (as well as tyre break and engine wear). This will be further elucidated in Section 3.4.5.

LDI-BB was highly correlated with offline AMS-BBOA ( $R^2=0.8283$ ), yet the LDI-BB concentrations were higher than BBOA from oAMS, especially in northern Switzerland (with an LDI-BB:AMS-BBOA ratio between 1.24 and 4.21 for the different sites). A possible reason is also here the mixing of secondary components into LDI-BB when comparing to primary BBOA from oAMS. However, we cannot exclude that this effect is due to different ~~ionization efficiencies~~rRFs of the LDI-MS for different compound classes.

The identified secondary components, LMW-OA ( $R^2=0.4745$ ) and bio-OA ( $R^2=0.6162$ ), also correlate with the corresponding OOA factors from the oAMS analysis, WOOA and SOOA, yet the correlation coefficients are smaller than

for LDI-BB and BBOA. For these factors, differences between the two methodologies could be related to differences in the PMF performance or to differences in the response factors for different components in the LDI-MS (the LMW-OA: WOOA ratio is between 0.67 and 1.01 and the bio-OA : SOOA ratio between 0.42 and 0.4).

5 At some of the sites analyzed with LDI-MS in this study, OA was monitored in previous years with state-of-the-art online aerosol mass spectrometry (AMS or ACSM). Earlier campaigns with quantitative online AMS analyses show a higher SOA contribution at those sites: e.g. in Zurich, July 2005 (66% vs 25% for LDI-MS summer) and December 2005 (55% vs 14% for LDI-MS winter), in Roveredo (close to San Vittore), in March 2005 (53% vs 46% for LDI-MS summer), and in December 2005 (43% vs 11% for LDI-MS winter), in Payerne, in July 2005 (94% vs 47% for LDI-MS summer) and in  
10 December 2005 (71% vs 20% for LDI-MS winter) (Lanz et al., 2010). Canonaco et al. (2013) presented source apportionment results for the winter of 2011 and 2012 for Zurich (PM<sub>1</sub> ACSM), also with higher SOA contributions (71%) than the LDI-MS in 2013 (20%).

Overall, the source apportionment results based on LDI-MS data provide source separations with similar temporal behaviors  
15 as offline AMS and online AMS and ACSM analysis, yet seem to overestimate combustion related primary particle sources and underestimate secondary OA. However, we assume for this comparison that all factors give an equal response at a given concentration, i.e. the relative response factor ( $rRF_{LDI}$ ) of all factors to be 1. The relative contribution of a factor  $k$  to the total signal observed with the LDI-MS ( $r_{i,k,LDI}$ ) for a specific measurement ( $i$ ) depends on the factor concentration ( $g_{i,k,LDI}$ ) as well as the sum of all factors ( $\sum_k^p(g_{i,k,LDI})$ ) separated for the LDI-MS data. Assuming that differences between the LDI-  
20 PMF and the AMS-PMF arise solely from different  $rRF_{LDI}$ s for different factors,  $r_{i,k,LDI}$  is a function of the AMS factor concentrations ( $g_{i,k,AMS}$ ) and the  $rRF_{LDI}$  of this factor:

$$r_{i,k,LDI} = \frac{g_{i,k,LDI}}{\sum_k^p(g_{i,k,LDI})} = \frac{g_{i,k,AMS} * rRF_{k,LDI}}{\sum_k^p(g_{i,k,AMS} * rRF_{k,LDI})} \quad (4)$$

25 In order to determine  $rRF_{LDI}$  for the LDI-MS several strong assumptions are required: (1) the sum of traffic1 and traffic2 represents HOA, (2) the sum of WBeff, WBeff1 and WBeff2 represents BBOA, (3) LMW-OA represents WOOA, (4) bio-OA represents SOOA and (5) the AMS factors for which there is an LDI-MS equivalent, (1) - (4), are the only contributors to OA. In scenario 1, all the above assumption are considered to be true. For scenario 1, we estimate  $rRF_{LDI-BB}$  as 1.84,  $rRF_{LDI-LMW-OA}$  as 0.29 and  $rRF_{LDI-bio-OA}$  as 0.40 (using LDI-traffic as reference factor,  $rRF_{LDI-traffic} = 1.00$ ). The LDI-  
30 MS factor concentrations corrected using  $rRF_{LDI}$  show a close relation to the uncorrected LDI factor concentrations ( $R^2_{LDI-traffic}=0.78$ ,  $R^2_{LDI-BB}=0.94$ ,  $R^2_{LDI-LMW-OA}=0.80$ ,  $R^2_{LDI-bio-OA}=0.85$ , Fig. 8).

We tested the sensitivity of the  $rRF_{LDI}$  estimates to the assumptions (1) and (2) being wrong (scenario 2) and only assumption (5) being wrong (scenario 3). In scenario 2, we alter scenario 1 by comparing the sum of traffic1 and bio-OA to  
35 SOOA and the sum of BBeff2 and LMW-OA to WOOA when computing  $rRF_{LDI}$ . For scenario 2, LDI-MS factor concentrations corrected using  $rRF_{LDI}$  show also a close relation to uncorrected LDI-MS factor concentrations ( $R^2_{LDI-traffic2}=0.96$ ,  $R^2_{BBeff+BBineff1}=1.00$ ,  $R^2_{LDI-BBeff2+LMW-OA}=0.99$ ,  $R^2_{traffic1+bio-OA}=0.96$ ). In scenario 3, we alter scenario 1 by considering also AMS factors without an equivalent in the LDI-MS PMF. Therefore, we compare LDI-traffic to the sum of HOA, COA, and SC-OA. Under these conditions LDI-MS factor concentrations corrected using  $rRF_{LDI}$  show a close relation  
40 to uncorrected LDI-MS factor concentrations ( $R^2_{LDI-traffic}=0.75$ ,  $R^2_{LDI-BB}=0.90$ ,  $R^2_{LMW-OA}=0.74$ ,  $R^2_{bio-OA}=0.85$ ). The differences in  $rRF_{LDI}$  between scenarios 1, 2, and 3 highlight the uncertainties caused by plausible violations of the underlying assumption when determining  $rRF_{LDI}$ . Overall the estimates  $rRF_{LDI}$  are highly uncertain and their accurate determination

[needs to be the focus of future work. Given the high uncertainty of  \$rRF\_{LDI}\$  and the good correlation between  \$rRF\_{LDI}\$  corrected and uncorrected LDI-MS factor concentrations, we present uncorrected results without considering  \$rRF\_{LDI}\$ .](#)

#### 5 3.4.4 Uncertainty of PMF results

The uncertainty estimate,  $\sigma_{tot}$ , includes both the statistical uncertainty ( $\sigma_{bs}$ ) and the uncertainty arising from the intraday variability of the measurements ( $\sigma_{intrad}$ , see Fig. 89, Section 2.3.2 and SI). In order to assess the impact of the intraday variability consideration ( $\sigma_{intrad}$ ) in estimating the uncertainty besides the statistical uncertainty obtained from the bootstrapping approach ( $\sigma_{bs}$ ), we compared the ratios  $\sigma_{bs}/\sigma_{tot}$  for the different factors: for traffic1 the ratio was 0.83, for traffic2 0.75, for BBeff 0.86, for BBineff1 0.80, for BBineff2 0.72, for LMW-OA 0.68, and for bio-OA 0.68. Thus, it was important to propagate  $\sigma_{intrad}$ . The relative uncertainties for the median factor concentrations ranged within 0.15 (for traffic1), 0.16 (bio-OA), 0.17 (LMW-OA), 0.18 (BBineff2), 0.20 (traffic2), 0.22 (BBeff), and 0.28 (BBineff1). Unlike the relative error of  $\sigma_{tot}$  of traffic1 (0.63 at the 10<sup>th</sup> percentile concentration and 0.13 at the 90<sup>th</sup> percentile concentration), traffic2 (0.68 and 0.11), BBeff (0.55 and 0.18), BBineff1 (0.47 and 0.20), BBineff2 (0.37 and 0.14), and LMW-OA (0.27 and 0.16) the relative error of bio-OA (0.20 and 0.15) only depended weakly on the factor concentration.

Throughout the measurement campaign, subsets of filters were analyzed repeatedly in order to assess the repeatability over longer time periods (see details in Supplement, Fig. S10). Most factors did not show significant changes of the attributed concentration during the measurement campaign. However, BBeff showed decreasing and LMW-OA increasing concentrations as a function of the measurement time. This could suggest an uncertain separation of these 2 factors. However, for these long time delays the variability increased and only few samples were repeated with such long time delays. Furthermore, the intra-day variance largely explains the total variance (traffic1 97%, traffic2 94%, BBeff 85%, BBineff1 89%, BBineff2 82%, LMW-OA 79%, bio-OA 97%, details in Supplement Fig. S10).

#### 25 3.4.5 Factor variabilities and contributions

The time series of all factors are shown for the 9 sites in Fig. 910 as relative contributions to OA and summarized as yearly averages in Table 4Tab. 2. In the yearly average for all sites, traffic1 contributes 7%, traffic2 12%, BBeff 7%, BBineff1 17%, BBineff2 32%, LMW-OA 21%, and bio-OA 5% to OM as measured by LDI-MS.

At the southern Alpine Swiss sites, the wood burning influenced categories (WB=BBeff+BBineff1+ BBineff2) contribute more than at the northern sites (70% vs 50%). Moreover, in winter WB explains 81% and 61% of OM in the south and the north, respectively. This difference is mostly caused by the higher relative contribution of BBineff1 (34% in the south vs 14% in the north) since BBineff2 (28% and 30%) and BBeff (7% and 5%) do not show strong geographical differences. The ratio BBineff1/BB (Fig. 40b11b) shows enhanced contributions of BBineff1 at the southern Alpine sites, especially during high pollution episodes in winter. This suggests different wood burning regimes in the 2 regions, as already discussed above.

After scaling of the mass spectra to OM, absolute factor concentrations have still to be interpreted with caution. One reason for this is that the relative response factors of the sources/factors are not known, another is that certain species (in analogy to the cooking-aerosol sample) may not be detected by LDI-MS if externally mixed. In support of the quantitative interpretation of our results, we note that the sum of the wood burning related LDI-MS factors correlates well with the offline AMS counterpart ( $R^2_{LDI,AMS}=0.82$ , described in Section 3.4.3). On the yearly average 24% of the measured OM is apportioned to

secondary OA. In summer, a bigger fraction is attributed to secondary OA (35% compared to 16% in winter). 35% (summer) and 1% (winter) of the secondary OA is attributed to biogenic sources. For some samples from the same period in Magadino, also the fossil and non-fossil content of OC and EC was determined (Vlachou et al., in prep) based on the method of Zhang et al. (2012). Vlachou et al. (in prep.) observed increased  $OC_f/EC_f$  ratios in summer which suggests other fossil POA sources in summer than in winter, or secondary formation of fossil OA. The higher traffic1 / eBC<sub>tr</sub> and traffic1/EC<sub>f</sub> ratios in summer are in agreement with enhanced  $OC_f/EC_f$  ratios (Fig. 4a, 11a). However, a part of traffic can also be mixed into BB leading to an underestimation of the traffic concentrations in winter. Overall this suggests that traffic1 represents aged or secondary traffic OA.

#### 4. Summary and conclusion

In this study, we developed/advanced a novel/known method for the chemical characterization of particulate matter collected on quartz-fiber filters by LDI-MS. We and applied the method to 819 samples. The method included the use of silver nitrate for  $m/z$  calibration and the automated peak integration of the mass spectra at unit-mass resolution. The benefit of LDI-MS measurements for the chemical characterization and a better understanding of the sources contributing to the ambient PM<sub>10</sub> was assessed at nine sites in central Europe throughout the entire year 2013.

Wood combustion smog chamber experiments revealed an influence of the burning conditions on the mass spectral signature. Tunnel samples used as a reference for traffic related emissions show mass spectral signatures distinctly different from wood combustion. Key  $m/z$ 's identified in the wood burning and traffic signatures showed links to expected markers as e.g., levoglucosan and NO<sub>x</sub>, respectively. The ambient mass spectral information was further used for source apportionment by PMF. Thereby, the influence of efficient and inefficient wood burning was separated. The extracted wood burning emissions correlated with the results from offline AMS source apportionment. Other components are more difficult to compare quantitatively because of different source separations in PMF as well as differences in the relative response factors ( $rRF$ ) of OA components.  $rRF$  determined in this study are uncertain and, therefore, not used for correcting the LDI-PMF results. The influence of traffic emissions was represented by 2 factors. One of these could clearly be linked to BC-related traffic (eBC<sub>tr</sub>) and NO<sub>x</sub>, and thus to primary emissions. The other, when normalized to eBC<sub>tr</sub>, showed a similar behavior as  $OC_f/EC_f$ , and was therefore attributed to aged/secondary traffic OA. A factor was attributed to biogenic SOA based on its concentration exponentially increasing with temperature. Another OA factor was characterized by low-molecular-weight ions and was correlated with NH<sub>4</sub><sup>+</sup> and was attributed to SOA from an unknown source.

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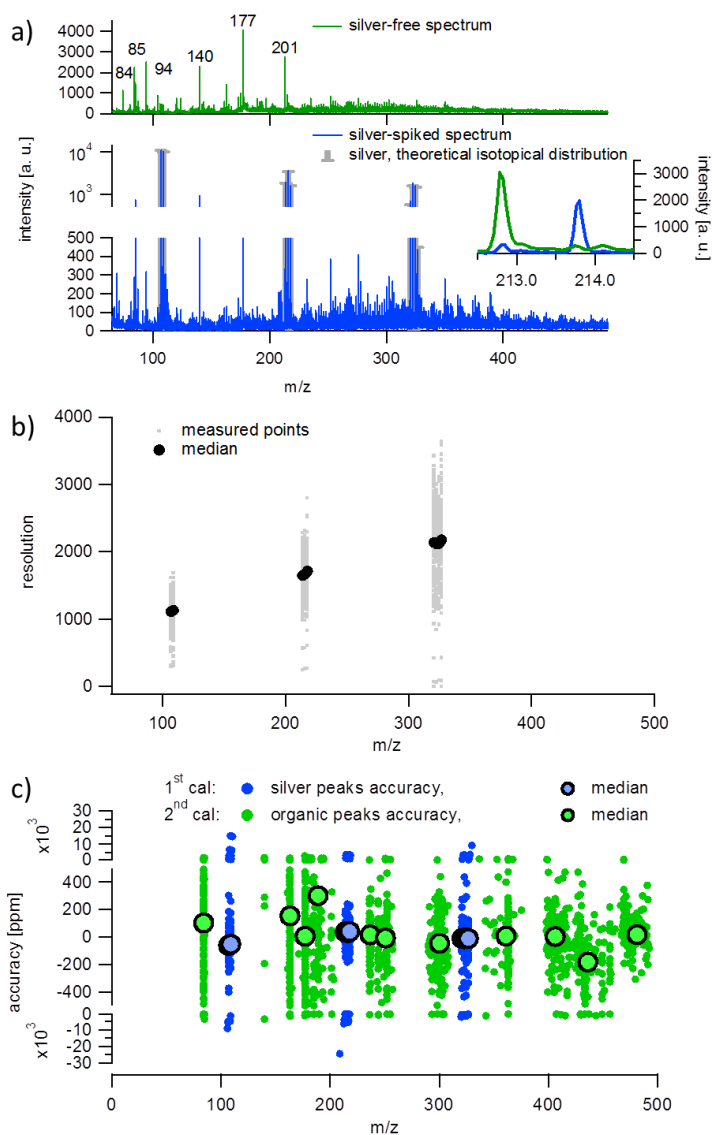


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**Figure 1:**  $m/z$  calibration of LDI-MS analysis of aerosol collected on quartz-fiber filter: a) example of calibrated silver (Ag-spec, blue) and no-silver-containing average mass spectra (noAg-spec, green) of a filter sample. The insert in a) displays a zoom-in of the Ag-spec and noAg-spec. b) operational resolution determined based on silver-mono-, di-, and trimer. c)  $m/z$  calibration accuracy for both steps of the  $m/z$  calibration.

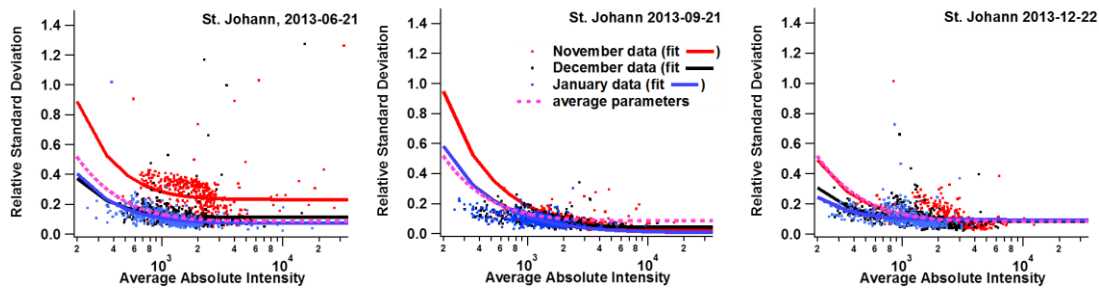


Figure 2: Error model parametrization based on 3 samples (Basel, 2013-06-21, 2013-09-21, 2013-12-22) measured on 3 instances with each time 10 repeats on 1 sample holder.

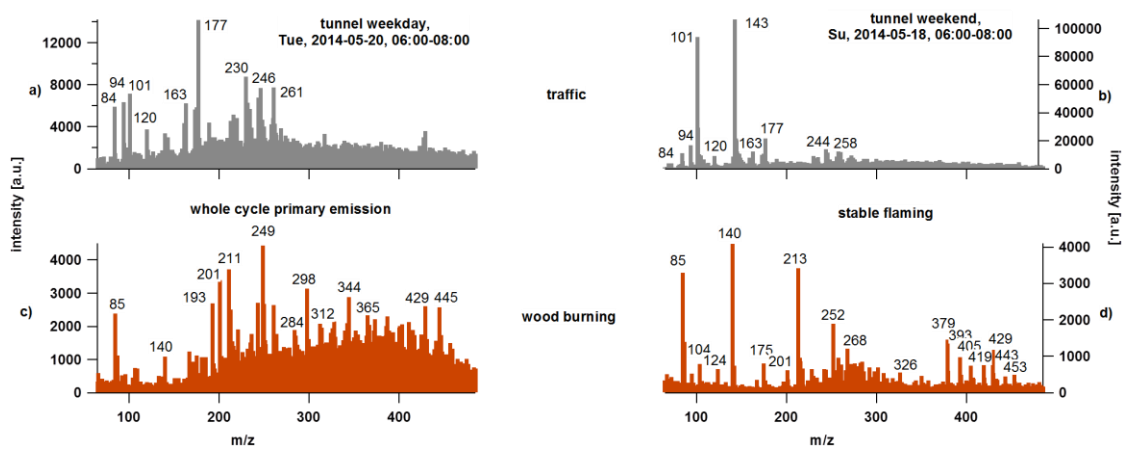


Figure 3: samples representing different combustion sources a) and b) traffic emissions from samples collected in Islisberg tunnel (exit, Wettswil, Switzerland), and c) primary wood burning using whole cycle emissions d) and stable flaming phase emissions d). In absence of a measurable mass spectrum, no spectrum for cooking emissions is displayed.

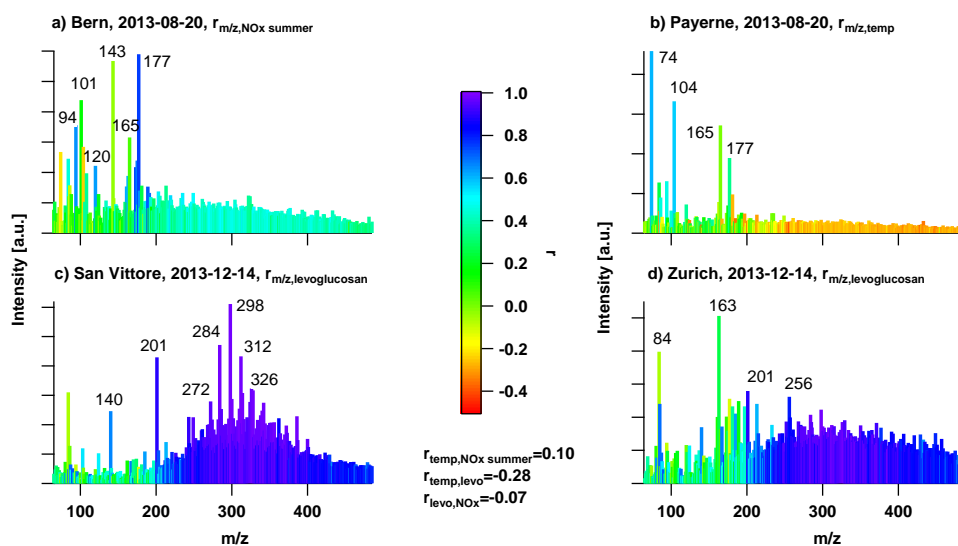
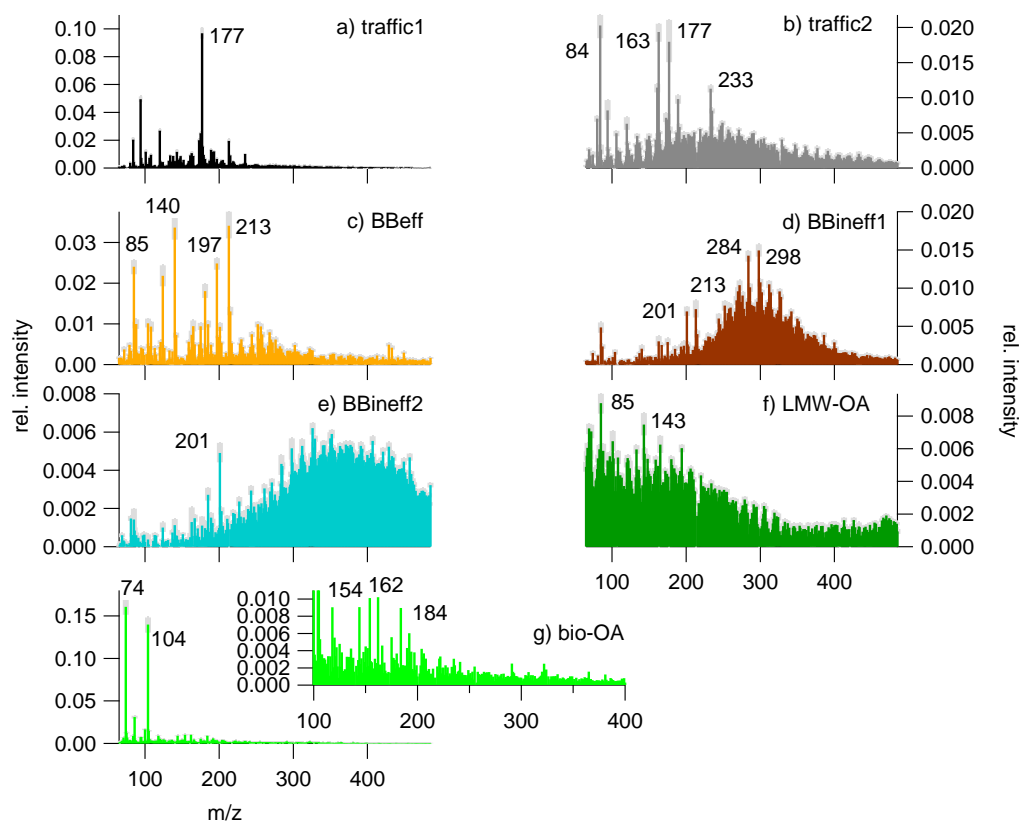


Figure 4: LDI-MS mass spectra from summer and winter samples (dates given in the legend) from a traffic-influenced (Bern), a rural (Payerne), a wood burning-influenced (San Vittore) and an urban background (Zurich) site. Spectra are color-coded with the correlation coefficient ( $r$ ) between the  $m/z$ s and a specific environmental parameter for the whole dataset: for a) the correlation with  $NO_x$  during summer, for b) the correlation with temperature, for c) and d) the correlation with levoglucosan.



**Figure 5:** PMF factor profiles (colored sticks) and their uncertainty (grey shaded areas, variability among PMF runs): a) traffic1, b) traffic2, c) efficient wood burning (BBeff), d) inefficient wood burning (BBineff1), e) inefficient wood burning 2 (BBineff2), f) lower molecular weight OA (LMW-OA), and g) biogenic OA (bio-OA).



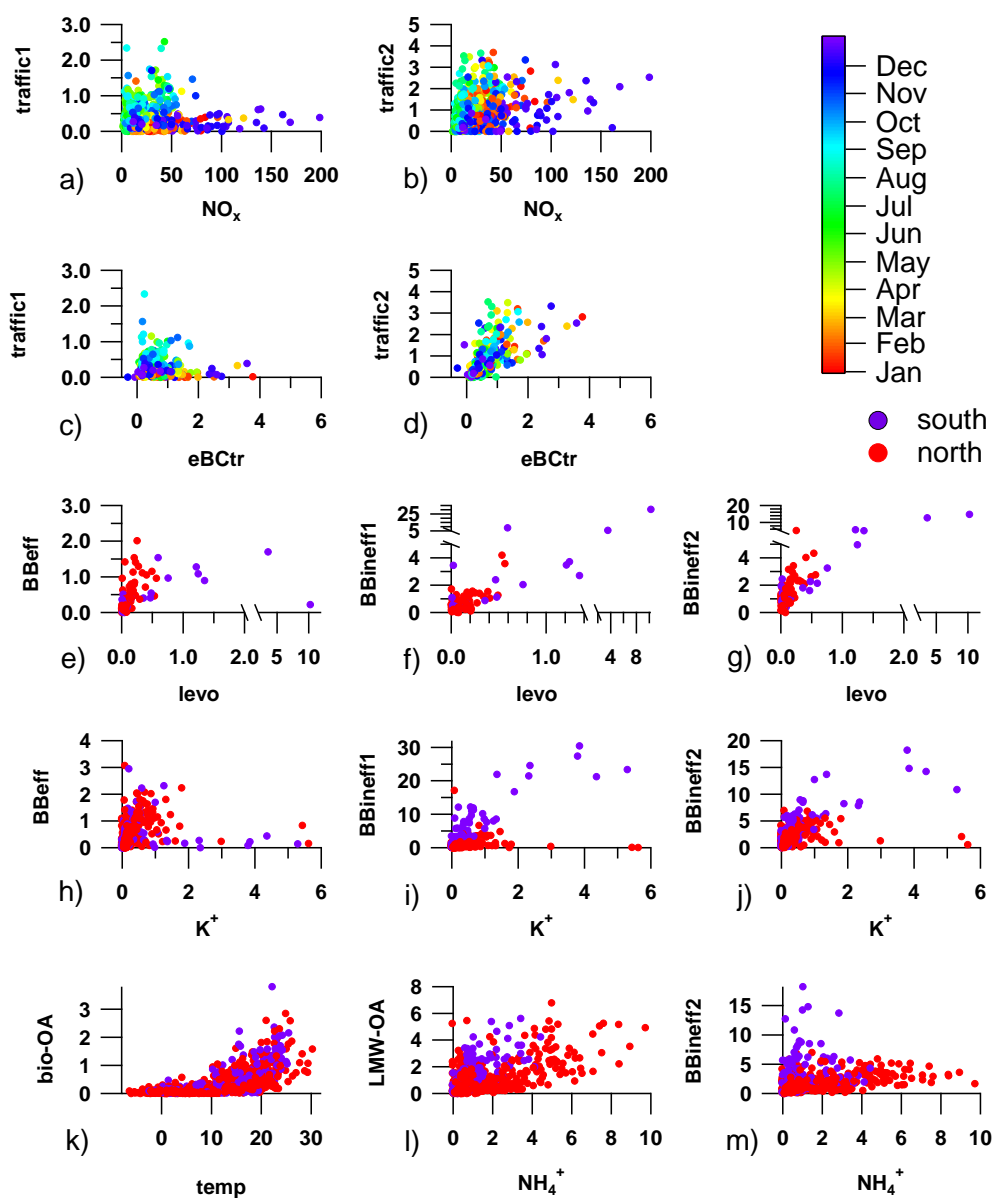
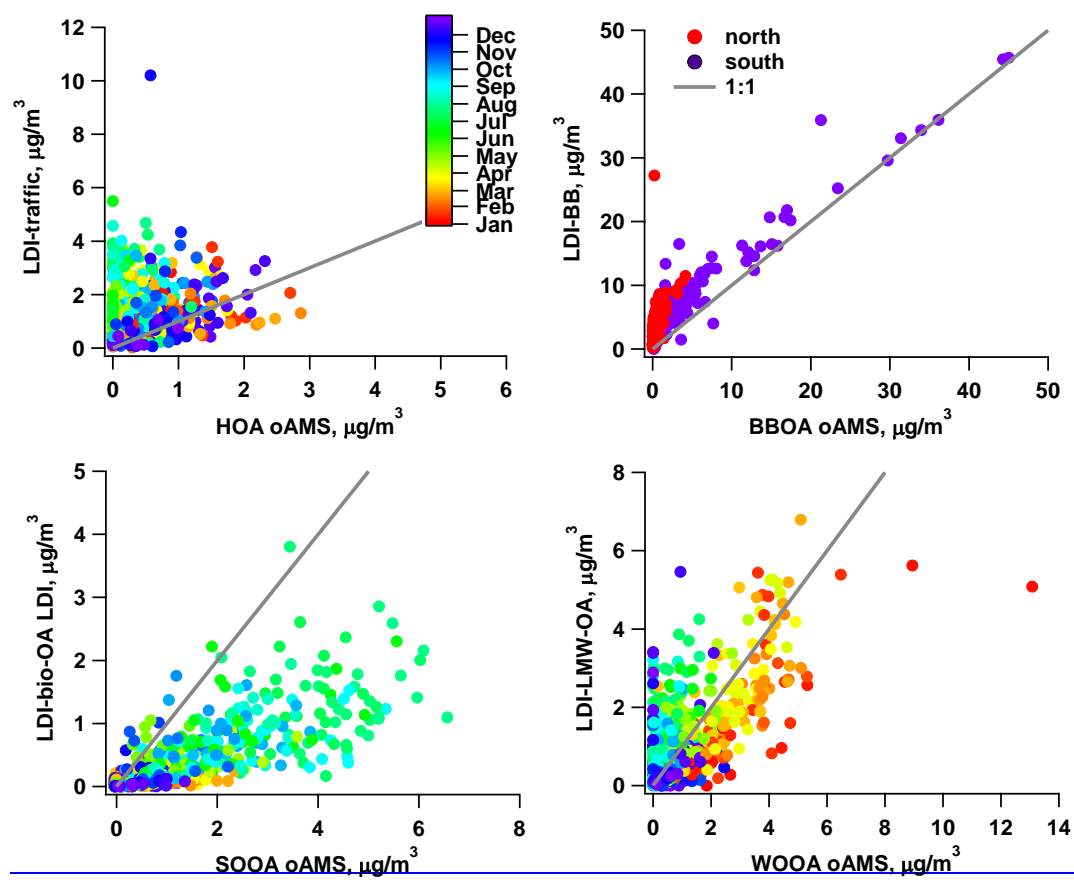


Figure 6: Scatterplots between factor-time series and respective markers (traffic1, traffic2, BB<sub>eff</sub>, BB<sub>ineff1</sub>, BB<sub>ineff2</sub>, bio-OA, LMW-OA, eBC<sub>tr</sub>, levoglucosan, potassium, ammonium are displayed in  $\mu\text{g}/\text{m}^3$ , NO<sub>x</sub> in ppm, and temperature in  $^{\circ}\text{C}$ ).



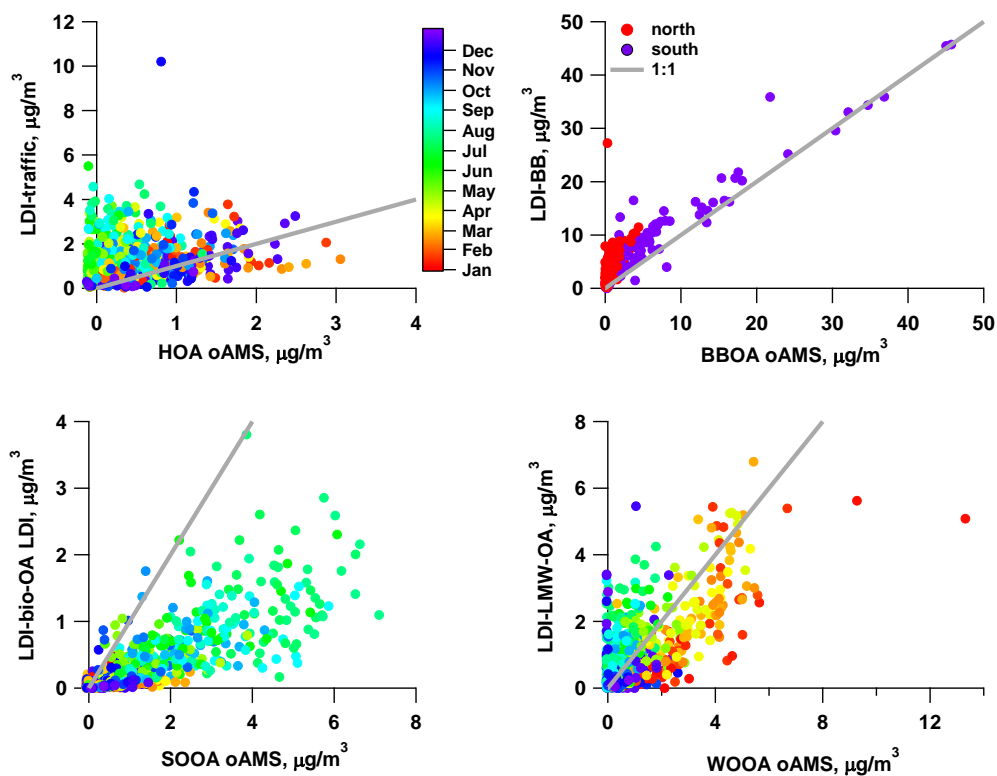


Figure 7: Comparison of LDI-MS to reference offline AMS source apportionment results for the sum of traffic related factors (LDI-traffic,  $R^2_{\text{LDI oAMS}}=0.4504$ ), sum of wood burning related factors (LDI-BB,  $R^2_{\text{LDI oAMS}}=0.8283$ ), bio-OA ( $R^2_{\text{LDI oAMS}}=0.6162$ ), and LMW-OA ( $R^2_{\text{LDI oAMS}}=0.4745$ ).

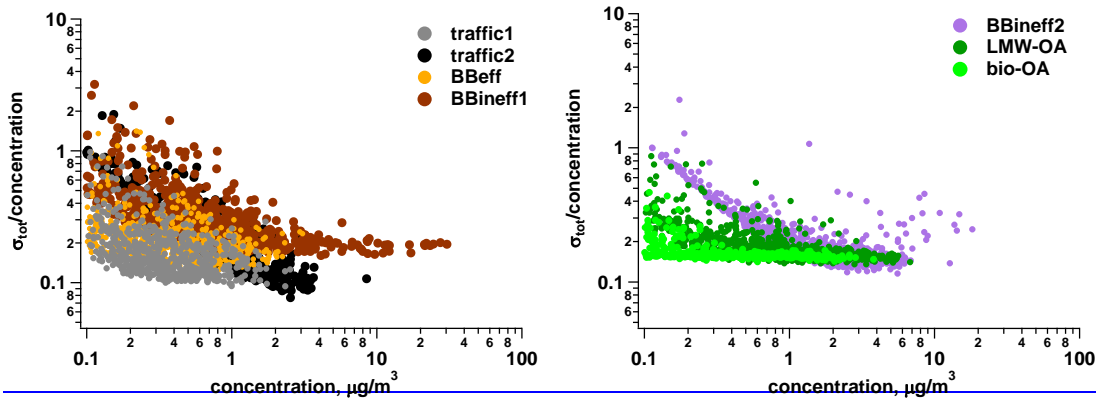


Table 1: relative response factors (*rRF*) for LDI-MS analyses for 3 different scenarios.

scenario 1	scenario 2	scenario 3
$rRF_{LDI-traffic} = 1.00 \pm 0.00$	$rRF_{LDI-traffic2} = 1.00 \pm 0.00$	$rRF_{LDI-traffic} = 1.00 \pm 0.00$
$rRF_{LDI-BB} = 1.84 \pm 0.06$	$rRF_{LDI-BB_{eff}+LDI-BB_{ineff1}} = 1.26 \pm 0.06$	$rRF_{LDI-BB} = 5.67 \pm 0.13$
$rRF_{LDI-LMW-OA} = 0.29 \pm 0.02$	$rRF_{LDI-LMW-OA+LDI-BB_{ineff2}} = 1.67 \pm 0.07$	$rRF_{LDI-LMW-OA} = 0.30 \pm 0.02$
$rRF_{LDI-bio-OA} = 0.14 \pm 0.01$	$rRF_{LDI-bio-OA+LDI-traffic1} = 0.39 \pm 0.02$	$rRF_{LDI-bio-OA} = 0.65 \pm 0.02$

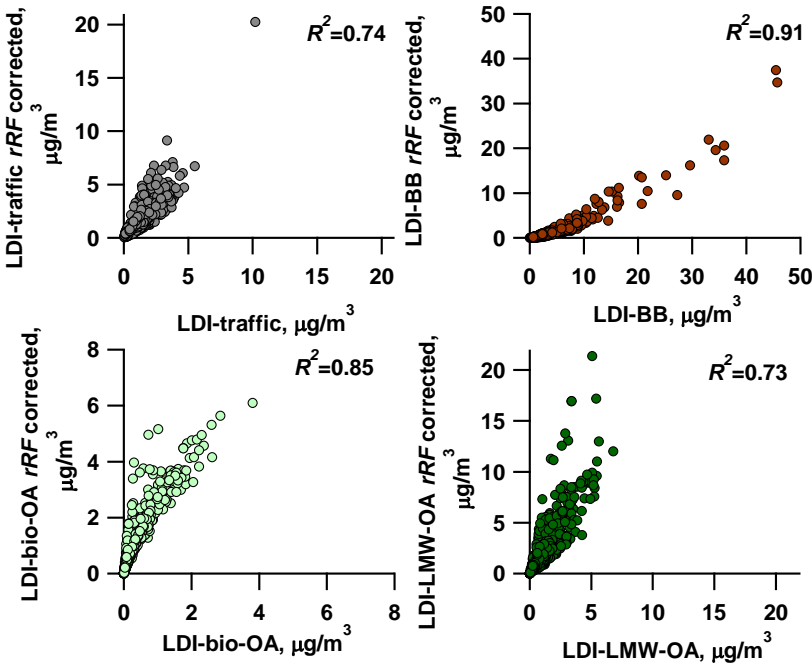
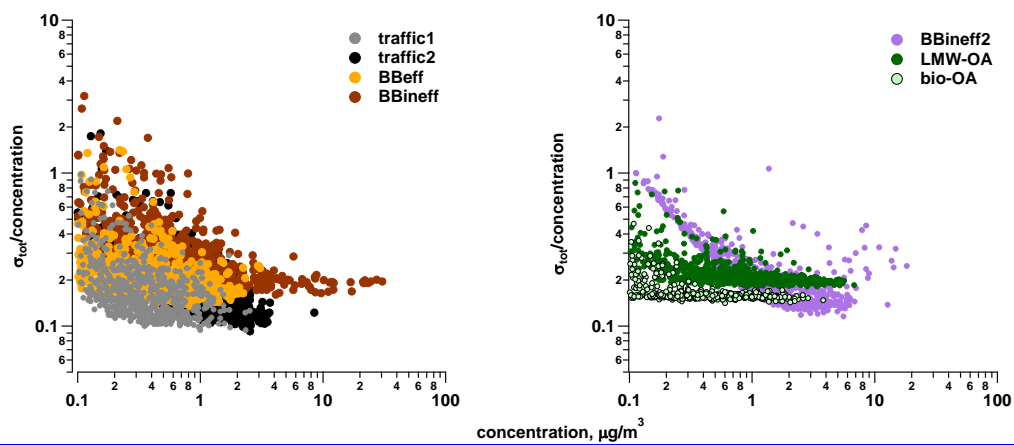


Figure 8: Comparison of LDI-MS factor concentrations corrected using relative response factors (*rRF*) to uncorrected LDI factor concentrations (Scenario 1).



**Figure 9:** Relative  $\sigma_{tot}$  for the different PMF factors as a function of the factor concentration.

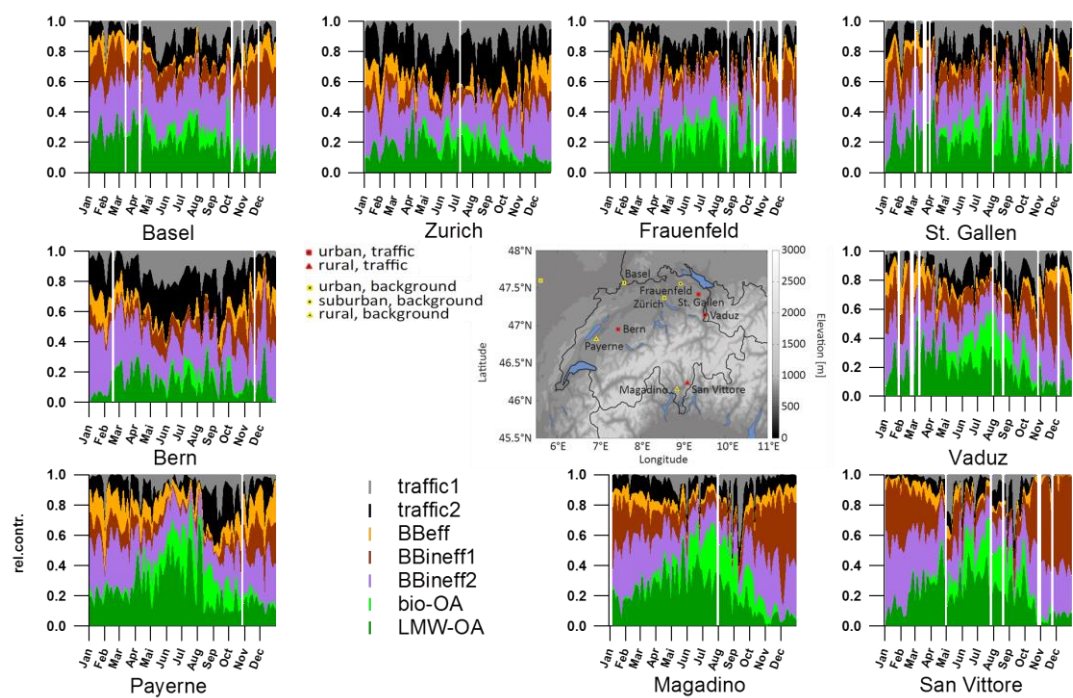
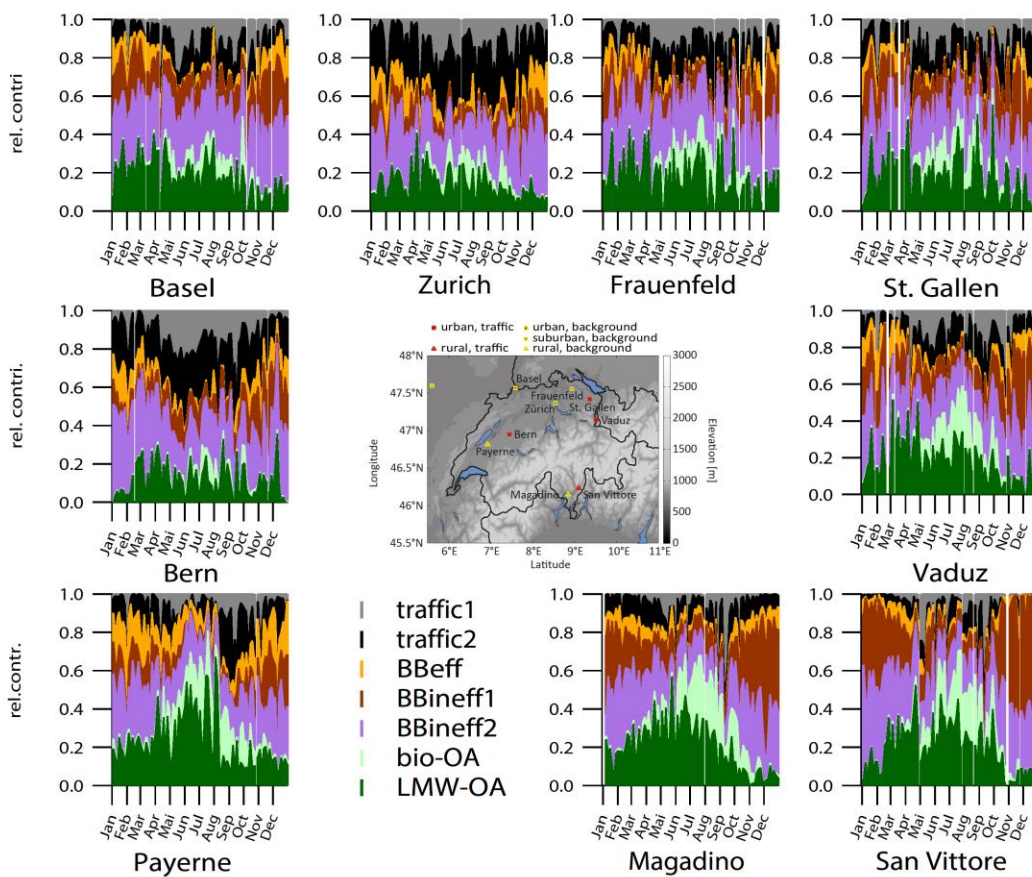


Figure 9



**Figure 10:** Relative factor time series of 7 identified factors for all nine sites in study area.

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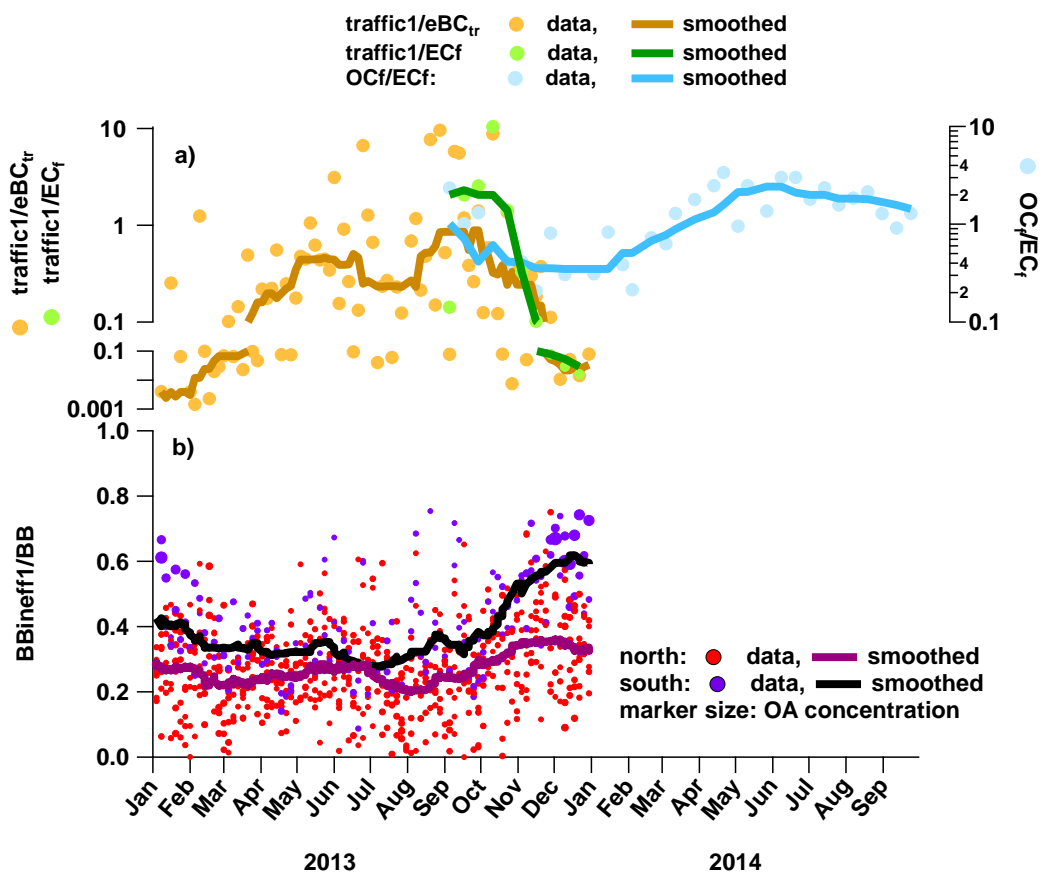


Figure 1140: a) time series of traffic1 normalized to  $\text{eBC}_{tr}$  and  $\text{EC}_f$  in comparison to  $\text{OC}_f/\text{EC}_f$  in Magadino, b) influence of inefficient wood burning emissions ( $\text{BBineff1}$ ) in comparison to the sum of wood burning influenced factors ( $\text{BB}=\text{BBineff1}+\text{BBineff2}+\text{BBeff}$ ) for the entire datasets.

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Table 24: yearly averages of the relative factor contributions and NOx concentrations.

yearly average, %	traffic1	traffic 2	BBeff	BBineff1	BBineff2	bio-OA	LMW-OA	NOx, ppb
<del>Bern</del>	<del>611</del>	<del>1224</del>	7	<del>1713</del>	<del>3230</del>	<del>53</del>	<del>2112</del>	<del>48</del>
<del>Zurich</del>	<del>149</del>	<del>2427</del>	<del>78</del>	<del>139</del>	<del>3027</del>	<del>35</del>	<del>1215</del>	<del>24</del>
<del>St. Gallen</del>	8	16	64	1317	3027	7	2021	24
<del>Basel</del>	86	1612	47	17	2732	75	21	21
<del>Frauenfeld</del>	38	1116	76	2713	2730	87	1720	21
<del>Magadino</del>	57	15	12	97	2520	1427	237	20
<del>Vaduz</del>	75	15	12	79	2025	2711	723	209
<del>Payerne</del>	23	411	37	4027	3227	58	1417	20
<del>S. Vittore</del>	92	274	83	940	2732	5	1514	18
<del>Zurich</del>								

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