

1 **New particle formation in the sulfuric acid-dimethylamine-water system:**
2 **Reevaluation of CLOUD chamber measurements and comparison to an**
3 **aerosol nucleation and growth model**
4
5

6 Andreas Kürten¹, Chenxi Li², Federico Bianchi³, Joachim Curtius¹, António Dias⁴, Neil M.
7 Donahue⁵, Jonathan Duplissy³, Richard C. Flagan⁶, Jani Hakala³, Tuija Jokinen³, Jasper
8 Kirkby^{1,7}, Markku Kulmala³, Ari Laaksonen⁸, Katrianne Lehtipalo^{3,9}, Vladimir Makhmutov¹⁰,
9 Antti Onnela⁷, Matti P. Rissanen³, Mario Simon¹, Mikko Sipilä³, Yuri Stozhkov¹⁰, Jasmin
10 Tröstl⁹, Penglin Ye^{5,11}, and Peter H. McMurry²

11
12 ¹Institute for Atmospheric and Environmental Sciences, Goethe University Frankfurt, 60438
13 Frankfurt am Main, Germany.

14 ²Department of Mechanical Engineering, University of Minnesota, 111 Church St. SE,
15 Minneapolis, MN 55455, USA.

16 ³Department of Physics, University of Helsinki, FI-00014 Helsinki, Finland.

17 ⁴SIM, University of Lisbon, 1849-016 Lisbon, Portugal.

18 ⁵Center for Atmospheric Particle Studies, Carnegie Mellon University, Pittsburgh,
19 Pennsylvania 15213, USA.

20 ⁶Division of Chemistry and Chemical Engineering, California Institute of Technology,
21 Pasadena, California 91125, USA.

22 ⁷CERN, CH-1211 Geneva, Switzerland.

23 ⁸Finnish Meteorological Institute, FI-00101 Helsinki, Finland.

24 ⁹Laboratory of Atmospheric Chemistry, Paul Scherrer Institute, 5232 Villigen PSI, Switzerland.

25 ¹⁰Solar and Cosmic Ray Research Laboratory, Lebedev Physical Institute, 119991 Moscow,
26 Russia.

27 ¹¹Aerodyne Research Inc., Billerica, Massachusetts 01821, USA.

28
29 Correspondence to: Andreas Kürten (kuerten@iau.uni-frankfurt.de)

30 **Abstract**

31

32 A recent CLOUD (Cosmics Leaving OUtdoor Droplets) chamber study showed that sulfuric
33 acid and dimethylamine produce new aerosols very efficiently, and yield particle formation
34 rates that are compatible with boundary layer observations. These previously published new
35 particle formation (NPF) rates are re-analyzed in the present study with an advanced method.
36 The results show that the NPF rates at 1.7 nm are more than a factor of 10 faster than previously
37 published due to earlier approximations in correcting particle measurements made at larger
38 detection threshold. The revised NPF rates agree almost perfectly with calculated rates from a
39 kinetic aerosol model at different sizes (1.7 nm and 4.3 nm mobility diameter). In addition,
40 modeled and measured size distributions show good agreement over a wide range (up to ca. 30
41 nm). Furthermore, the aerosol model is modified such that evaporation rates for some clusters
42 can be taken into account; these evaporation rates were previously published from a flow tube
43 study. Using this model, the findings from the present study and the flow tube experiment can
44 be brought into good agreement for the high base to acid ratios (~100) relevant for this study.
45 This confirms that nucleation proceeds at rates that are compatible with collision-controlled
46 (a.k.a. kinetically-controlled) new particle formation for the conditions during the CLOUD7
47 experiment (278 K, 38% RH, sulfuric acid concentration between 1×10^6 and 3×10^7 cm^{-3} and
48 dimethylamine mixing ratio of ~40 pptv, i.e., 1×10^9 cm^{-3}).

49 1. INTRODUCTION

50

51 The formation of new particles by gas-to-particle conversion (nucleation or new particle
52 formation, NPF) is important for a variety of atmospheric processes and for human health.

53 It has been shown in numerous studies that sulfuric acid (H_2SO_4) is often associated with
54 NPF (Weber et al., 1997; Kulmala et al., 2004; Fiedler et al., 2005; Kuang et al., 2008; Kirkby
55 et al., 2011) and indeed it can explain some of the observed particle formation together with
56 water vapor for neutral (uncharged) and ion-induced conditions when temperatures are low,
57 e.g., in the free troposphere (Lee et al., 2003; Lovejoy et al., 2004; Duplissy et al., 2016; Ehrhart
58 et al., 2016; Dunne et al., 2016). However, at least one additional stabilizing compound is
59 required in order to explain boundary layer nucleation at warm temperatures. Acid-base
60 nucleation, which involves a ternary compound, e.g., ammonia, besides sulfuric acid and water,
61 can lead to much higher NPF rates compared to the binary system (Weber et al., 1998; Ball et
62 al., 1999; Kürten et al., 2016a). Nevertheless, for most conditions close to the surface, the
63 concentrations of H_2SO_4 and NH_3 are too low, or temperatures are too high, to allow significant
64 ternary nucleation of these compounds (Kirkby et al., 2011; Kürten et al., 2016a). However, the
65 substitution of ammonia by amines, e.g., dimethylamine ($(\text{CH}_3)_2\text{NH}$), leads to NPF rates that
66 can explain the atmospheric observations over a wide range of sulfuric acid concentrations,
67 even when the amine mixing ratios are in the low pptv-range (Kurtén et al., 2008; Nadykto et
68 al., 2011; Ortega et al., 2012; Chen et al., 2012; Almeida et al., 2013; Glasoe et al., 2015). A
69 recent study even showed that NPF is collision-controlled, i.e., that it proceeds at the maximum
70 possible speed (Rao and McMurry, 1989), when amine mixing ratios are above ~ 20 pptv (5×10^8
71 cm^{-3}), and sulfuric acid concentrations are between $1 \times 10^6 \text{ cm}^{-3}$ and $3 \times 10^7 \text{ cm}^{-3}$ at 278 K and
72 38% RH (Kürten et al., 2014). Indications that NPF can be collision-limited were reported more
73 than 30 years ago based on the analysis of chamber nucleation experiments (McMurry, 1980),
74 although the involvement of amines, which were probably present as a contaminant during
75 those experiments, was not considered. Indications that atmospheric nucleation might occur by
76 a collision-limited process have also been previously presented (Weber et al., 1996). Despite
77 the strong evidence that sulfuric acid-amine nucleation is very efficient, it has rarely been
78 observed in the atmosphere. Only one study has so far reported sulfuric acid-amine nucleation
79 (Zhao et al., 2011) despite amine mixing ratios of up to tens of pptv at some sites (Yu and Lee,
80 2012; You et al., 2014; Freshour et al., 2014; Yao et al., 2016). A global modelling study of
81 sulfuric acid-amine nucleation has been carried out so far (Bergman et al., 2015) applying a
82 nucleation parametrization based on the measurements of Almeida et al. (2013) and Glasoe et
83 al. (2015).

84 Atmospheric boundary layer nucleation can also be explained by the existence of highly-
85 oxygenated organic molecules (Crouse et al., 2013; Ehn et al., 2014), e.g., from α -pinene.
86 These highly-oxygenated molecules have been found to nucleate efficiently in a chamber study
87 even without the involvement of sulfuric acid, especially when ions take part in the nucleation
88 process (Kirkby et al., 2016).

89 Even though oxidized organics seem to be globally important for NPF (Jokinen et al., 2015;
90 Gordon et al., 2016; Dunne et al., 2016), the formation of new particles by sulfuric acid and
91 amines should still be considered because sulfuric acid-amine nucleation rates exceed those
92 from oxidized organics as soon as the concentrations of the precursor gases (sulfuric acid and

93 amines) are high enough (Berndt et al., 2014). Therefore, at least locally or regionally, i.e., close
94 to sources, amines should be relevant.

95 In this study, we reanalyze data from CLOUD (Cosmics Leaving OUtdoor Droplets)
96 chamber experiments conducted at CERN during October/November 2012 (CLOUD7
97 campaign). New particle formation rates as a function of the sulfuric acid concentration from
98 CLOUD7 were previously published (Almeida et al., 2013). However, these data are re-
99 analyzed in the present study using an advanced method that takes into account the effect of
100 self-coagulation in the estimation of new particle formation rates (Kürten et al., 2015a). The re-
101 analyzed data and NPF rates obtained from Scanning Mobility Particle Sizer (SMPS)
102 measurements are compared to results from a kinetic aerosol model. Modeling is also used for
103 a comparison between results from a flow tube study (Jen et al., 2016a) and CLOUD.

104 The reanalyzed data cover sulfuric acid concentrations from ca. 1×10^6 to 3×10^7 cm^{-3} , which
105 fall into the range for most observations of atmospheric boundary layer new particle formation
106 events (e.g. Kulmala et al., 2013). The dimethylamine mixing ratio for most of the data shown
107 in this study is ~ 40 pptv (1×10^9 cm^{-3}), which is within the rather wide range of observations
108 (0.1 to 157 pptv, i.e., 2.5×10^6 to 4×10^9 cm^{-3}) for C2-amines to which dimethylamine belongs to
109 (Yao et al., 2016).

110
111

112 2. METHODS

113

114 2.1 CLOUD experiment and instruments

115

116 The CLOUD (Cosmics Leaving OUtdoor Droplets) experiment at CERN was designed to
117 investigate nucleation and growth of aerosol particles in chemically diverse systems.
118 Additionally, the influence of ions on new particle formation (NPF) and growth can be studied
119 inside the 26.1 m^3 electro-polished stainless steel chamber (Kirkby et al., 2011). For the
120 experiments discussed in this paper, NPF is initiated by illuminating the air inside the chamber
121 with UV light by means of a fiber-optic system (Kupc et al., 2011), which produces sulfuric
122 acid (H_2SO_4) photolytically from reactions involving O_3 , H_2O , SO_2 and O_2 . Diluted
123 dimethylamine and sulfur dioxide are taken from gas bottles; inside the chamber, these trace
124 gases mix with clean synthetic air (i.e., O_2 and N_2 with a ratio of 21:79 from evaporated
125 cryogenic liquids). To ensure homogenous conditions, the air is mixed with magnetically driven
126 fans installed at the top and bottom of the chamber (Voigtländer et al., 2012). A thermal housing
127 controls the chamber temperature to 278.15 K within several hundredths of a degree. The
128 temperature was not varied for the experiments relevant for this study. The relative humidity
129 was kept constant at 38% by humidifying a fraction of the inflowing air with a humidification
130 system (Duplissy et al., 2016). In order to keep the pressure inside the chamber at 1.005 bar,
131 the air that is taken by the instruments has to be continuously replenished. Therefore, a flow of
132 150 l/min of the humidified air is continuously supplied to the chamber. For the sulfuric acid,
133 dimethylamine and water system, ions do not have a strong enhancing effect on the nucleation
134 rates for most conditions (Almeida et al., 2013); therefore, we do not distinguish between the
135 neutral and charged pathway in such runs.

136 A suite of instruments is connected to the CLOUD chamber to measure particles, ions,
 137 clusters and gas concentrations. A summary of these instruments is provided elsewhere (Kirkby
 138 et al., 2011; Duplissy et al., 2016). For this study, measured sulfuric acid and particle
 139 concentrations are relevant. A Chemical Ionization-Atmospheric Pressure interface-Time Of
 140 Flight Mass Spectrometer (CI-APi-TOF) was employed to measure sulfuric acid and its neutral
 141 clusters in this study (Jokinen et al., 2012; Kürten et al., 2014). The particle concentrations
 142 originate from a scanning mobility particle sizer (SMPS, Wang and Flagan, 1990), which
 143 measured the particle size distribution between ~4 and ~80 nm. The SMPS uses a differential
 144 mobility analyzer built by the Paul Scherrer Institute; it includes a Kr⁸⁵ charger to bring the
 145 particles into a charge equilibrium before they are classified. The retrieval of the particle size
 146 distributions requires corrections for the charging and the transmission efficiency, which were
 147 performed according to the literature (Wiedensohler and Fissan, 1988; Karlsson and
 148 Martinsson, 2003). The mixing ratio of dimethylamine was determined by ion chromatography
 149 with a detection limit of 0.2 to 1 pptv (5×10^6 to 2.5×10^7 cm⁻³) at a time resolution between 70
 150 and 210 minutes (Praplan et al., 2012; Simon et al., 2016).

151

152 2.2 Calculation of particle formation rates

153

154 Particle formation rates J (cm⁻³ s⁻¹) are calculated from the measured size distributions (assumed
 155 to consist of n bins). For the size bin with the index m , the rate at which particles with a diameter
 156 equal or larger than d_m are formed can be calculated according to Kürten et al., 2015a:

157

$$158 J_{\geq m} = \frac{dN_{\geq m}}{dt} + \sum_{i=m}^n (k_{w,i} \cdot N_i) + k_{dil} \cdot N_{\geq m} + \sum_{i=m}^n (\sum_{j=i}^n S_{i,j} \cdot K_{i,j} \cdot N_j \cdot N_i). \quad (1)$$

159

160 This equation takes into account the time derivative of the number density of all particles for
 161 which $d_p \geq d_m$, i.e., $N_{\geq m}$, and corrects for the effects of wall loss (size dependent wall loss rates
 162 $k_{w,i}$), dilution (dilution rate k_{dil}), and coagulation (collision frequency function $K_{i,j}$), where N_i
 163 and N_j are the particle number densities in different size bins. The rate of losses to the chamber
 164 walls can be expressed by Crump and Seinfeld, 1981:

165

$$166 k_w(d_p) = C_w \cdot \sqrt{D(d_p)}, \quad (2)$$

167

168 where $D(d_p)$ is the diffusivity of a particle of diameter d_p , which is given by the Stokes-Einstein
 169 relation (Hinds, 1999),

170

$$171 D(d_p) = \frac{k_B \cdot T \cdot C_C}{3 \cdot \pi \cdot \eta \cdot d_p}, \quad (3)$$

172

173 where k_b , T , η , are the Boltzmann constant, the temperature, and the gas viscosity, respectively.
 174 The Cunningham slip correction factor, C_C , is a function of the particle Knudsen number, $Kn =$
 175 $2\lambda/d_p$, and λ is the mean-free-path of the gas molecules. The empirically derived proportionality
 176 coefficient, C_w , depends upon the chamber dimensions and on the intensity of turbulent mixing.

177 The rate of loss of sulfuric acid to the chamber walls is generally used to characterize C_w . The
 178 diffusivity of sulfuric acid is $0.0732 \text{ cm}^2 \text{ s}^{-1}$ at 278 K and 38% RH (Hanson and Eisele, 2000).
 179 The measured life time, determined from the decay of sulfuric acid when the UV light is turned
 180 off, was 554 s (wall loss rate 0.00181 s^{-1}), with the experimentally determined diffusivity this
 181 yields a factor C_w of $0.00667 \text{ cm}^{-1} \text{ s}^{-0.5}$. However, in this study diffusivities were calculated
 182 according to equation (3), so the calculated monomer diffusivity (for a monomer with a density
 183 of 1470 kg m^{-3} and a molecular weight of $0.143 \text{ kg mol}^{-1}$, see section 2.4) required a different
 184 scaling, resulting in a value of $C_w = 0.00542 \text{ cm}^{-1} \text{ s}^{-0.5}$ that was used throughout this study.

185 Dilution is taken into account by a loss rate that is independent of size and equals $k_{\text{dil}} =$
 186 $9.6 \times 10^{-5} \text{ s}^{-1}$. Correcting for particle-particle collisions requires the calculation of the collision
 187 frequency function. We used the method from Chan and Mozurkewich (2001). This method
 188 includes the effect of enhanced collision rates through van der Waals forces. A value of
 189 $6.4 \times 10^{-20} \text{ J}$ was used for the Hamaker constant (Hamaker, 1937), leading to a maximum
 190 enhancement factor of ~ 2.3 for the smallest clusters, relative to the collision rate in the absence
 191 of van der Waals forces. The factor of 2.3 has previously been shown to give good agreement
 192 between measured and modeled cluster and particle concentrations for the chemical system of
 193 sulfuric acid and dimethylamine (Kürten et al., 2014; Lehtipalo et al., 2016). In order to consider
 194 the collisions of particles in the same size bin, a scaling factor $s_{i,j}$ is used in equation (1), which
 195 is 0.5 when $i = j$ and 1 otherwise.

196

197 **2.3 Reconstruction method**

198

199 Recently a new method was introduced, that makes it possible to retrieve new particle formation
 200 rates at sizes below the threshold of the instrument used to determine the particle number
 201 density. This method is capable of considering the effect of self-coagulation (Kürten et al.,
 202 2015a). It requires introducing new size bins below the threshold of the SMPS (termed d_{p2} in
 203 the following; d_{p2} corresponds to the index $m = 1$). The method starts by calculating the number
 204 density in the first newly introduced smaller size bin (index $m = 0$, diameter $d_{p2} - dd_p$):

205

$$206 \quad N_{m-1} = (d_{p,m} - d_{p,m-1}) \cdot \frac{J_{\geq m}}{GR_{m-1}} \approx dd_p \cdot \frac{J_{\geq m}}{GR}. \quad (4)$$

207

208 Here, the particle growth rate GR (nm s^{-1}) needs to be used as well as the difference between
 209 two adjacent size bins (dd_p). Once the number density in the newly introduced bin is known
 210 this information can be used to calculate J_{m-1} . In the further steps, the numbers N_{m-2} and J_{m-2} are
 211 calculated and so on. In this way, the size distribution can be extrapolated towards smaller and
 212 smaller sizes in a stepwise process until eventually reaching the diameter d_{p1} .

213 The method has so far only been tested against simulated data but not against measured size
 214 distributions (Kürten et al., 2015a). In this study the smallest measured SMPS diameter is $d_{p2} =$
 215 4.3 nm ; 26 new size bins with $dd_p = 0.1 \text{ nm}$ were introduced and this enabled the calculation of
 216 the NPF rates at $d_{p1} = 1.7 \text{ nm}$ in the smallest size bin. This size was chosen since previously
 217 published particle formation rates from the CLOUD experiment were reported for this diameter
 218 (e.g. Kirkby et al., 2011; Almeida et al., 2013; Riccobono et al., 2014).

219 The method introduced here explicitly takes into account losses that occur between particles
 220 with d_{p1} and d_{p2} (self-coagulation). These losses have not been taken into account by Almeida

221 et al. (2013). Almeida et al. (2013) derived $J_{3,2\text{nm}}$ from CPC and SMPS measurements by
 222 including the corrections for wall loss, dilution and coagulation above 3.2 nm (see also Kürten
 223 et al., 2016a). However, the extrapolation to 1.7 nm was made by using the Kerminen and
 224 Kulmala equation (Kerminen and Kulmala, 2002), which does not include the effect of self-
 225 coagulation. For the system of sulfuric acid and dimethylamine, where a significant fraction of
 226 particles reside in the small size range, this process is, however, important.

228 **2.4 Kinetic new particle formation and growth model**

229
 230 The measured particle formation rates are compared to modeled formation rates assuming
 231 collision-limited particle formation, i.e., all clusters are not allowed to evaporate. McMurry
 232 (1980) was the first to show that number concentrations and size distributions of particles
 233 formed photochemically from SO₂ in chamber experiments (Clark and Whitby, 1975) are
 234 consistent with collision-controlled nucleation; results from updated versions of this model
 235 have recently been presented (Kürten et al., 2014; McMurry and Li, 2017). The model used
 236 here has been described previously (Kürten et al., 2014; Kürten et al., 2015a, Kürten et al.
 237 2015b) but only brief introductions were reported; therefore, more details are provided in the
 238 following.

239 As outlined in Kürten et al. (2014), collision-controlled new particle formation accurately
 240 described the measured cluster distributions for the sulfuric acid-dimethylamine system up to
 241 the pentamer (cluster containing five sulfuric acid molecules). In this model, it was assumed
 242 that the clusters consist of “monomeric” building blocks, each containing one dimethylamine
 243 and one sulfuric acid molecule. Evidence that this 1:1-ratio between base and acid is
 244 approximately maintained for the small clusters was presented from neutral and charged cluster
 245 measurements (Almeida et al., 2013; Kürten et al., 2014; Bianchi et al., 2014; Glasoe et al.,
 246 2015). The molecular weight was, therefore, chosen as 0.143 kg mol⁻¹ (sum of sulfuric acid
 247 with 0.098 kg mol⁻¹ and dimethylamine with 0.045 kg mol⁻¹), and the density as 1470 kg m⁻³
 248 (Qiu and Zhang, 2012).

249 During the reported experiments (CLOUD7 in fall 2012), dimethylamine was always present
 250 at mixing ratios above ca. 20 pptv (5×10^8 cm⁻³). Dimethylamine (DMA) was supplied from a
 251 certified gas bottle and diluted with synthetic air before it was introduced into the chamber to
 252 achieve the desired mixing ratios. Sulfuric acid was generated in situ from the reactions between
 253 SO₂ and OH whenever the UV light was turned on (see section 2.1). Since the UV light intensity
 254 and the gas concentrations were kept constant throughout each individual experiment, it is
 255 justified to assume a constant monomer production rate P_1 . The equation describing the
 256 temporal development of the monomer concentration, N_1 , is

$$257 \frac{dN_1}{dt} = P_1 - (k_{1,w} + k_{dil} + \sum_{j=1}^{N_{max}} K_{1,j} \cdot N_j) \cdot N_1 \quad (5)$$

258 and, for the clusters containing two or more sulfuric acid molecules ($k \geq 2$),

$$261 \frac{dN_k}{dt} = \frac{1}{2} \cdot \sum_{i+j=k} K_{i,j} \cdot N_i \cdot N_j - (k_{w,k} + k_{dil} + \sum_{j=1}^N K_{k,j} \cdot N_j) \cdot N_k. \quad (6)$$

264 The same loss mechanisms (wall loss, dilution and coagulation) as for the calculation of the
 265 particle formation rates (section 2.2) are considered when modeling the cluster concentrations.
 266 In this study, the particle size distribution was calculated from the monomer up to a diameter
 267 of ~84 nm, which corresponds to the upper size limit of the SMPS used in CLOUD7. Tracking
 268 each individual cluster/particle up to this large size would be computationally too demanding,
 269 so the size distribution was divided into so-called molecular size bins (tracking each individual
 270 cluster), and geometric size bins, where the mid-point diameters of two neighboring size bins
 271 differ by a constant factor. The number of molecular size bins was set to 400 (which results in
 272 a diameter of ~5 nm for the largest molecular bin), while the number of geometric size bins was
 273 set to 190 with a geometric factor of 1.015 (maximum diameter of the last bin is 83.7 nm). The
 274 treatment of the geometric size bins was similar to the molecular bins, except that the collision
 275 products were distributed between the two closest size bins. Two smaller particles with
 276 diameters $d_{p,i}$ and $d_{p,j}$ generate a cluster with size

$$277 \quad d_{p,x} = (d_{p,i}^3 + d_{p,j}^3)^{1/3}. \quad (7)$$

278 If it is assumed that the collision product falls into the size range covered by the geometric bins,
 281 its diameter will be between two size bins $d_{p,k}$ and $d_{p,k+1}$. The production rate of particles with
 282 diameter $d_{p,x}$ is

$$283 \quad P_x = s_{i,j} \cdot K_{i,j} \cdot N_i \cdot N_j. \quad (8)$$

284 For the geometric size range, the resulting particles are distributed between the two bins to
 285 conserve mass, i.e.,

$$286 \quad P_k = \left(\frac{d_{p,k+1}^3 - d_{p,x}^3}{d_{p,k+1}^3 - d_{p,k}^3} \right) \cdot P_x, \quad (9a)$$

$$287 \quad P_{k+1} = \left(1 - \frac{d_{p,k+1}^3 - d_{p,x}^3}{d_{p,k+1}^3 - d_{p,k}^3} \right) \cdot P_x. \quad (9b)$$

288
 289
 290
 291
 292 When the collision product falls into the molecular size bin regime the calculation is
 293 straightforward because the diameter of the product agrees exactly with a molecular bin and
 294 does not need to be distributed between two bins (see the production term in equation (6)). In
 295 case the collision products exceed the largest bin diameter, the product is entirely assigned to
 296 the largest bin, while taking into account the scaling such that the total mass is conserved.

297 In the model, no free parameter is used as the concentration of monomers is constrained by
 298 the measurements. Therefore, the production rate P_1 is adjusted such that the resulting monomer
 299 concentration in the model matches the measured sulfuric acid concentration. The model is used
 300 to simulate the experiments for a duration of 10,000 s with a time resolution of 1 s. For the
 301 small clusters and particles this leads to a steady-state between production and loss; therefore,
 302 the resulting concentrations are essentially time-independent.

303 The model introduced here was compared with the model described in McMurry and Li
 304 (2017) and yielded almost indistinguishable results for several scenarios when the same input
 305 parameters were used. We take this as an indication that both models correctly describe

306 collision-controlled nucleation, especially since the models were independently developed and
307 do not share the same code. The model in this paper is based on defining size bins according to
308 their diameter, while the model by McMurry and Li (2017) uses particle volume.

309

310 **2.5 Nucleation and growth model involving selected evaporation rates**

311

312 Measured cluster concentrations for the sulfuric acid-dimethylamine system from flow tube
313 experiments indicated that finite evaporation rates exist for some clusters (Jen et al., 2014; Jen
314 et al., 2016a). This was supported by the observation that diamines can yield even higher
315 formation rates than amines for some conditions (Jen et al., 2016b). Within the flow tube
316 experiments dimethylamine was mixed into a gas flow containing a known amount of sulfuric
317 acid monomers. The products, i.e., the sulfuric acid-dimethylamine clusters were measured
318 after a short reaction time (≤ 20 s) with a chemical ionization mass spectrometer. From the
319 measured signals, the cluster evaporation rates were retrieved from model calculations (Jen et
320 al., 2016a). The main differences to the CLOUD study lie within the much shorter reaction time
321 (20 s vs. steady state in CLOUD) and in the much wider range of base to acid ratios used by
322 Jen et al. (2016a, 2016b). This allowed them to retrieve even relatively slow evaporation rates
323 for the sulfuric acid-dimethylamine clusters. The measured cluster/particle concentrations
324 increased with increasing base to acid ratio, eventually approaching a plateau at a
325 dimethylamine to acid ratio of ~ 1 . Therefore, the high dimethylamine to acid ratio used in the
326 CLOUD7 experiment (~ 100) can probably explain why our NPF rates are compatible with
327 collision-controlled nucleation.

328 However, this was further tested by incorporating the evaporation rates from Jen et al.
329 (2016a) in our model. For this purpose, the model described in section 2.4 was modified in a
330 way that allows retrieving the cluster concentrations of the monomer, dimer, trimer and tetramer
331 as a function of their dimethylamine content (see Appendix A). The abbreviation A_xB_y denotes
332 the concentration of a cluster containing x sulfuric acid ($x = 1$ for the monomer) and y base
333 (dimethylamine) molecules. It is assumed that $x \geq y$ for all clusters, i.e., the number of bases is
334 always smaller or equal to the number of acid molecules. The reported cluster concentrations
335 (Fig. 3) refer to the number of acid molecules in the cluster, i.e., $N_1 = A_1 + A_1B_1$, $N_2 = A_2B_1 +$
336 A_2B_2 and $N_3 = A_3B_1 + A_3B_2 + A_3B_3$.

337 The evaporation rates considered are $k_{e,A_1B_1} = 0.1 \text{ s}^{-1}$, $k_{e,A_3B_1} = 1 \text{ s}^{-1}$, $k_{e,A_3B_2} = 1 \text{ s}^{-1}$ (Jen et al.,
338 2016a). Jen et al. (2016a) suggested that the formation of stable tetramers requires at least two
339 base molecules. In this case the evaporation rate of k_{e,A_4B_1} is infinity. In the model, this was
340 solved by not taking into account the formation of clusters A_4B_1 (from A_3B_1 and A_1) at all.
341 Further details about the modeling involving evaporation rates can be found in Appendix A and
342 in Table 1, which gives a summary over the different model studies.

343

344

345 **3. RESULTS**

346

347 **3.1 Comparison between Almeida et al. (2013) and SMPS derived NPF rates**

348

349 Using the model described in section 2.4, a comparison between the previously published NPF
350 rates from Almeida et al. (2013) and the modeled rates was performed. Almeida et al. (2013)
351 derived NPF rates for a particle mobility diameter of 1.7 nm. Using a density of 1470 kg m^{-3}
352 and a molecular weight of $0.143 \text{ kg mol}^{-1}$, it can be calculated that a spherical cluster containing
353 nine monomers (nonamer) has a geometric diameter of $\sim 1.4 \text{ nm}$, i.e., a mobility diameter of 1.7
354 nm (Ku and Fernandez de la Mora, 2009, see also Appendix A); therefore, the modeled nonamer
355 formation rates were used for the comparison.

356 Figure 1 shows the modeled formation rates at 1.7 nm and the Almeida et al. (2013) data as
357 a function of the sulfuric acid concentration (which is equivalent to the monomer concentration
358 in the model, see section 2.4, since it is assumed that all sulfuric acid is bound to DMA). It can
359 be seen that the modeled NPF rates are significantly higher. This indicates that the previously
360 published formation rates underestimate the true formation rates if sulfuric acid-dimethylamine
361 nucleation is indeed proceeding at the collision-limit. Previously published results indicated
362 that this is the case (Kürten et al., 2014; Lehtipalo et al., 2016); however, we will provide further
363 evidence that this assumption accurately describes the experiments in the present study and
364 provide an explanation why Almeida et al. (2013) underestimated the formation rates.

365 It should be noted that the displayed experimental $J_{1.7\text{nm}}$ values (open red triangles in Fig. 1)
366 are identical to the values from Almeida et al. (2013), while the sulfuric acid concentration has
367 been corrected. In Almeida et al. (2013) data were shown from CLOUD4 (spring 2011) and
368 CLOUD7 (fall 2012). For consistency, the sulfuric acid concentrations from the chemical
369 ionization mass spectrometer (Kürten et al., 2011) were used, as the CI-APi-TOF was not
370 available during CLOUD4. Especially during CLOUD7, the chemical ionization mass
371 spectrometer (CIMS) showed relatively high sulfuric acid concentrations even when no sulfuric
372 acid was produced from the UV light system inside the CLOUD chamber; no correction was
373 applied for this effect in Almeida et al. (2013). However, taking into account a subtraction of
374 this instrumental background (reaching sometimes values above $1 \times 10^6 \text{ cm}^{-3}$) leads to a
375 shallower slope for $J_{1.7\text{nm}}$ vs. sulfuric acid and brings the corrected CIMS values in a good
376 agreement with the sulfuric acid measured by the CI-APi-TOF. In the present study, the data
377 from the CI-APi-TOF were used. The slope for $J_{1.7\text{nm}}$ vs. sulfuric acid now yields a value of
378 close to 2, while the previously reported value was ~ 3.7 (Almeida et al., 2013). The higher
379 value resulted from the bias in the sulfuric acid concentration and the consideration of data
380 points at low sulfuric acid concentration, where new particle formation is significantly affected
381 by losses to the chamber walls, which tends to bias the slope towards higher values (Ehrhart
382 and Curtius, 2013).

383

384 **3.2 Comparison between NPF rates from the kinetic model and SMPS measurements**

385

386 The formation rates in Almeida et al. (2013) were calculated from measured particle number
387 densities with a condensation particle counter that has a lower cut-off diameter of $\sim 3 \text{ nm}$. The
388 derivation of particle formation rates at 1.7 nm therefore required an extrapolation to the smaller
389 diameter (Kerminen and Kulmala, 2002). With the available model, we are now, in principle,
390 able to calculate NPF rates for any particle diameter and compare the result to directly measured
391 rates. This was done for the SMPS size channel corresponding to a mobility diameter of 4.3 nm
392 ($J_{4.3\text{nm}}$) with the method described in section 2.2. Using the SMPS data has the advantage that

393 the size-dependent loss rates can be accurately taken into account, which is not possible when
394 only the total (non size-resolved) concentration from a condensation particle counter is
395 available. On the other hand, the smallest SMPS size channels need to be corrected by large
396 factors to account for losses and charging probability (section 2.1), which introduces
397 uncertainty.

398 The result for $J_{4.3\text{nm}}$ is shown in Figure 1 together with the modeled particle formation rates
399 for the same diameter. The agreement between modeled and measured NPF rates is very good
400 indicating that the collision-controlled model accurately describes 4.3 nm particle production
401 rates for these experiments. This is further evidence that particles are formed at the collision-
402 limit. However, it is also an indication that the Almeida et al. (2013) data underestimate the
403 NPF rates, which is further discussed in the following section.

404

405 **3.3 Reconstruction model results**

406

407 Recently, a new method was introduced, which allows the extrapolation of NPF rates
408 determined at a larger size (d_{p2}) to a smaller diameter (d_{p1}). The advantage of that method is
409 that the effect of cluster-cluster collisions (self-coagulation) can be accurately taken into
410 account (Kürten et al., 2015a). So far, the method has not been tested for measured particle size
411 distributions. However, the effect of cluster-cluster collisions should be largest in the case of
412 collision-controlled nucleation since it results in the highest possible cluster (particle)
413 concentrations for a given production rate of nucleating molecules. Therefore, the current data
414 set is ideal for testing the new method. It requires the measured growth rate as an input
415 parameter (equation (4)); this growth rate was derived from fitting a linear curve to the mode
416 diameter determined from the SMPS size distribution (Hirsikko et al., 2005). It was then used
417 as a constant (i.e., it was assumed that it is independent of size) for the full reconstruction of
418 the size distribution, in order to obtain a formation rate at 1.7 nm. The growth rate could only
419 be accurately determined for experiments with relatively high sulfuric acid concentration
420 (above $\sim 5 \times 10^6 \text{ cm}^{-3}$); therefore, the reconstruction method was only tested for these conditions
421 (Figure 1). The comparison with the modeled formation rates at the same size (1.7 nm) shows
422 that the reconstruction method yields quite accurate results, highlighting the importance of
423 cluster-cluster collisions in this chemical system. This explains why the Almeida et al. (2013)
424 data strongly underestimate the particle formation rates.

425 While the reconstruction method gives good results in the present study, it needs to be
426 mentioned that the errors for this method can become quite large. Small inaccuracies in the
427 growth rate, can be blown up to very large uncertainties due to the non-linear nature of the
428 method. This can be seen for some of the data points with large error bars in the positive
429 direction. The errors are calculated by repeating the reconstruction with growth rates $GR \pm dGR$,
430 where dGR ($\pm 20\%$) is the error from the fitted growth rate. Therefore, the accuracy of the
431 method strongly depends on good growth rate measurements, and relies on the assumption that
432 the growth rate does not change as a function of size. This seems to be a reasonable
433 approximation for collision-controlled nucleation under the present conditions (Kürten et al.,
434 2015a), but it could be different in other chemical systems.

435 The higher formation rates are also consistent with calculations from the ACDC
436 (Atmospheric Cluster Dynamics Code) model (McGrath et al., 2012) that were previously

437 published in Almeida et al. (2013). Figure 1 shows the rates calculated by the ACDC model
 438 (black lines). It should be noted that these values refer to a mobility diameter of 1.2 to 1.4 nm
 439 and therefore, somewhat higher rates are expected due to the smaller diameter compared to
 440 $J_{1.7nm}$. However, the agreement between the measured and predicted rates from ACDC are now
 441 in much better agreement than before.

442 Hanson et al. (2017) recently reported an expression for the calculation of particle formation
 443 rates as a function of the sulfuric acid concentration, dimethylamine concentration and
 444 temperature. According to their formula the formation rate of tetramers (mobility diameter of
 445 ~ 1.4 nm, see Appendix A) follows the expression

$$447 \quad J_{1.4nm} = \exp\left(-129 + \frac{16200 K}{T}\right) \cdot \left(\frac{N_1}{cm^{-3}}\right)^3 \cdot \left(\frac{DMA}{cm^{-3}}\right)^{1.5} \quad (10)$$

448
 449 The formation rates $J_{1.4nm}$ are shown in Fig. 1 (green line) for a DMA mixing ratio of 40 pptv
 450 (1×10^9 cm⁻³) and a temperature of 278 K. At the first glance, the agreement between the
 451 experimental CLOUD data and the ACDC simulation is remarkably good. However, one should
 452 note that Hanson et al. (2017) recommended to use their equation only for DMA between 2
 453 pptv (5×10^7 cm⁻³) and 16 pptv (4×10^8 cm⁻³) if sulfuric acid is present between 1×10^6 cm⁻³ and
 454 2×10^7 cm⁻³. Using the equation in this range avoids that the formation rates can exceed the
 455 kinetic limit. When using larger concentrations, the kinetic limit is eventually exceeded due to
 456 the power dependency of 3 regarding sulfuric acid and the 1.5 power dependency for DMA.
 457 Further comparison between equation (10) and the results from the present study are shown in
 458 Fig. 3 (lower panel).

459

460 **3.4 Size distribution comparison between model and SMPS**

461

462 Further comparison between modeled and measured data was performed for one experimental
 463 run (CLOUD7 run 1036.01) in which the particles were grown to sizes beyond 20 nm.
 464 Therefore, the time-dependent cluster/particle concentrations were modeled for a monomer
 465 production rate of 2.9×10^5 cm⁻³ s⁻¹, which results in a steady-state monomer concentration of
 466 1.07×10^7 cm⁻³ for the model; this is the same as the measured sulfuric acid concentration. The
 467 measured and modeled size distributions are shown in Fig. 2 (panels a, b and c) at four different
 468 times, i.e., at 1h, 2h, 4h and 6h after the start of the experiment. Given that there is no free
 469 parameter used in the model, the agreement between the base case simulation and the
 470 measurement is very good (Fig. 2a). For the earliest time shown (1h) the modeled
 471 concentrations overestimate the measured concentrations by up to 30%, whereas for the later
 472 times (≥ 4 h) the model underestimates the measured concentrations by up to 30%. It is unclear
 473 whether these discrepancies are due to SMPS measurement uncertainties, or if the model does
 474 not include or accurately describe all the relevant processes. If, for example, the SMPS would
 475 underestimate the concentrations of the smaller particles ($< ca. 15$ nm) and overestimate those
 476 of the larger particles, the observed difference between modeled and measured concentrations
 477 could also be explained.

478 A comparison between measured and modeled aerosol volume concentrations is shown
 479 in Fig. 2d. In order to enable direct comparison, the modeled size distribution was integrated

480 starting at 4.3 nm since the SMPS did not capture smaller particles. In the beginning of the
481 experiment the modeled aerosol volume is up to ~40% larger than the measured one, but,
482 towards the end of the experiment (ca. 4h after its start), the volumes agree quite well. Possibly
483 this is because the overestimated modeled particle number density at small diameters is
484 compensated by the underestimated particle concentration in the larger size range (see Fig. 2a).

485 This trend leads eventually to a slight underestimation of the aerosol volume by the model.

486 If one assumes that the SMPS is not responsible for the slight disagreement, then the
487 following conclusions can be drawn regarding the accuracy of the model. The particle growth
488 rate is well represented by the model given the good agreement between the positions of the
489 local maxima in the size distribution and the intersections between the size distributions and
490 the x -axis. This good agreement between measured and modeled growth rates has already been
491 demonstrated in Lehtipalo et al. (2016) for a particle diameter of 2 nm. The results shown here
492 indicate that no significant condensation of other trace gases contribute to the growth of
493 particles because, in this case, the measured particle size distributions would be shifted towards
494 larger diameters compared to the model.

495 The good agreement between model and measurement is also a confirmation of the effect of
496 van der Waals forces, when a Hamaker constant of 6.4×10^{-20} J is used, a value that has been
497 demonstrated previously to represent particle size distribution dynamics correctly (McMurry,
498 1980; Chan and Mozurkewich, 2001; Kürten et al., 2014; Lehtipalo et al., 2016). Regarding the
499 underestimation of the modeled size distribution for diameters ≥ 15 nm, one explanation could
500 be that the size-dependent particle loss rates in the CLOUD chamber are weaker than assumed
501 ($k_w \sim D^{0.5}$; see equation (2)). A weaker size dependence would lead to higher predicted particle
502 concentrations at larger sizes (Park et al., 2001). However, no evidence was found from the
503 existing CLOUD data that this is the case. Dedicated wall loss experiments could be performed
504 in the future to investigate this hypothesis further.

505 In order to test the model sensitivity to certain variations quantitatively further simulations
506 were performed (Fig. 2b and Fig. 2c). A variation of the steady-state sulfuric acid monomer
507 concentration by $\pm 20\%$ was achieved by using different monomer production rates for the high
508 sulfuric acid case ($P_1 = 4.17 \times 10^5 \text{ cm}^{-3} \text{ s}^{-1}$) and for the low sulfuric acid case ($P_1 = 2.01 \times 10^5$
509 $\text{cm}^{-3} \text{ s}^{-1}$, Fig. 2b). This rather small variation leads to significant mismatches between the
510 modeled and measured size distributions that is also found for the aerosol volumes (Fig. 2d).

511 Two further scenarios were tested with the model. First, the enhancement due to van der
512 Waals forces were turned off. This scenario results in significantly slower growth rates and the
513 modeled size distributions do not match the measured ones at all anymore (Fig. 2c); the same
514 is found when comparing modeled and measured aerosol volumes (Fig. 2d). Second, the aerosol
515 density and the molecular weight of the condensing “monomer” were changed. In the base-case
516 simulations (Fig. 2a), the density of dimethylammonium-bisulfate is 1470 kg m^{-3} and the
517 molecular weight is $0.143 \text{ kg mol}^{-1}$ because a one to one ratio between DMA and sulfuric acid
518 is assumed. Since full neutralization of sulfuric acid by DMA would require a 2:1-ratio between
519 base and acid, collision-controlled nucleation of $(\text{H}_2\text{SO}_4)((\text{CH}_3)_2\text{NH})_2$ “monomers” instead of
520 $(\text{H}_2\text{SO}_4)((\text{CH}_3)_2\text{NH})$ was tested. Therefore, the density was decreased by 6% to account for the
521 density change between dimethylammonium-bisulfate and dimethylammonium-sulfate (see Qiu and
522 Zhang, 2011) and the molecular weight was set to $0.188 \text{ kg mol}^{-1}$. As expected, the particle
523 growth is now slightly faster due to the additional volume added by the further DMA molecules

524 (Fig. 2c). However, the changes are rather small and the modeled size distributions move a little
525 further away from the measurements compared to the base case scenario (Fig. 2a).

526 Comparison between modeled and measured size distributions yielded similar results for
527 other experiments from CLOUD7. However, the experiment shown in Fig. 2 was carried out
528 over a relatively long time (6 h) at high sulfuric acid concentrations. Therefore, the particles
529 could grow to large diameters and the comparison between model and experiment covers a wide
530 size range.

531

532 **3.5 Sensitivity of cluster concentrations and NPF rates regarding DMA**

533

534 The data presented in the previous sections provide evidence that the new particle formation in
535 the sulfuric acid-dimethylamine system during CLOUD7 proceeds at rates that are consistent
536 with collision-controlled nucleation, in agreement with results for this data set obtained using
537 different approaches (Kürten et al., 2014; Lehtipalo et al., 2016). In this section, we compare
538 whether for CLOUD conditions the collision-controlled assumption is consistent with the Jen
539 et al. (2016a) results that showed that some clusters evaporate at the rates given in section 2.5
540 and Table. 1.

541 For the following discussion, both versions of the nucleation and growth model (section 2.4
542 and section 2.5) were used. Figure 3 shows a comparison between calculated cluster (dimer,
543 trimer, tetramer and pentamer) concentrations using collision-controlled nucleation (section
544 2.4) and the model described in section 2.5. When a DMA mixing ratio of 40 pptv ($1 \times 10^9 \text{ cm}^{-3}$)
545 is used (this was the average mixing ratio of DMA during the CLOUD7 experiments), there is
546 almost no difference between the two scenarios. This indicates that, under the CLOUD7
547 conditions, new particle formation proceeded at almost the same rates that result for collision-
548 controlled nucleation. Nevertheless, this does not imply that all cluster evaporation rates are
549 zero. The conditions are only such that, due to the high DMA mixing ratio, most of the clusters
550 (including the monomer) probably contain as many DMA molecules as sulfuric acid molecules;
551 this results in very stable cluster configurations (Ortega et al., 2012). When DMA mixing ratios
552 are low, most sulfuric acid clusters contain, however, only a small number of DMA molecules.
553 As these clusters can evaporate more rapidly, the overall formation rate is slowed down (Ortega
554 et al., 2012; Hanson et al., 2017). For low base to acid ratios, it can therefore matter whether a
555 cluster is stabilized by a dimethylamine, a diamine (Jen et al., 2016) or by both an amine and
556 an ammonia molecule (Glasoe et al., 2015). This can explain the more efficient NPF due to
557 diamines or the synergistic effects involving amines and ammonia at low base to acid ratios. At
558 high base to acid ratios, the differences in the effective evaporation rates become small (Jen et
559 al., 2016b).

560 The effect of the dimethylamine concentration on the cluster concentrations and on the
561 particle formation rate was further investigated. The lower panel of Fig. 3 shows that the cluster
562 concentrations and the NPF rate at 1.7 nm decrease with decreasing DMA levels. The figure
563 shows the concentrations and the NPF rate normalized by the results for the collision-limit. The
564 NPF rate drops by about a factor of three when DMA is reduced to $2.5 \times 10^7 \text{ cm}^{-3}$ (~ 1 pptv).
565 Below that level, the reduction in J and in the trimer, tetramer, and pentamer concentrations is
566 approximately linear with DMA. The dimer is less affected since, in the model, its evaporation
567 rates are set to zero while the evaporating trimers contribute to the dimer concentration. From

568 this perspective, very high particle formation rates should be observed even at DMA mixing
569 ratios around 1 pptv ($2.5 \times 10^7 \text{ cm}^{-3}$), which should be almost indistinguishable from rates
570 calculated for collision-controlled nucleation. Possibilities why such high rates have so far not
571 been observed are discussed in section 4.

572 For a comparison, the expected formation rates from equation (10) are shown in Fig. 3, lower
573 panel, by the grey line. The values were scaled similar to the simulated data by setting the value
574 for 40 pptv ($1 \times 10^9 \text{ cm}^{-3}$) to 1. Although this DMA mixing ratio is outside the range for which
575 the Hanson et al. (2017) formulation is recommended for (between $5 \times 10^7 \text{ cm}^{-3}$ and $4 \times 10^8 \text{ cm}^{-3}$),
576 from Fig. 1 it can be concluded that both, the Hanson et al. (2017) equation and the kinetic
577 model agree quite well at this DMA mixing ratio. The slope of J vs. DMA seems to be, however,
578 different in the relevant range of DMA ($5 \times 10^7 \text{ cm}^{-3}$ and $4 \times 10^8 \text{ cm}^{-3}$). This is due to the fact,
579 that the model predicts a steep slope (close to the value of 1.5 in equation (10)) only for much
580 lower DMA ($< 2.5 \times 10^6 \text{ cm}^{-3}$), for higher DMA the slope flattens out and reaches eventually a
581 plateau, when the value for collision-controlled nucleation is approached. This flattening of the
582 curve is not reflected in the simple formulation from Hanson et al. (2017). However, in contrast
583 to the three constant evaporation rates used in our modeling approach, Hanson et al. (2017)
584 used a more sophisticated nucleation scheme involving many different evaporation rates, not
585 only regarding sulfuric acid but also for dimethylamine. This more complex scheme was,
586 however, not implemented in our model.

587 Further experiments are required to derive accurate values for evaporation rates in the
588 sulfuric acid-dimethylamine system; these experiments should especially target DMA
589 concentrations with low base to acid ratios (< 10).

590
591

592 4. DISCUSSION

593

594 This study confirms the results derived in previous studies that new particle formation in the
595 sulfuric acid-dimethylamine-water system can proceed at or close to the collision-controlled
596 limit (Kürten et al., 2014; Lehtipalo et al., 2016). This is the case for sulfuric acid concentrations
597 between 1×10^6 and $3 \times 10^7 \text{ cm}^{-3}$ and dimethylamine mixing ratios around 40 pptv ($1 \times 10^9 \text{ cm}^{-3}$)
598 at 278 K and 38% RH. For these conditions particle formation rates and size distributions can
599 be reproduced with high accuracy by an aerosol model that assumes that particle growth is
600 exclusively due the irreversible addition of $\text{H}_2\text{SO}_4 \cdot (\text{CH}_3)_2\text{NH}$ “monomers” and coagulation.
601 Even when evaporation rates for the less stable clusters are introduced in the model (Jen et al.,
602 2016a) the resulting particle formation rates are effectively indistinguishable from the kinetic
603 model results for CLOUD7 conditions (i.e., at the high dimethylamine to acid ratio of ~ 100).
604 The fact that the measured particle size distribution can be reproduced with good accuracy
605 shows that neither water nor other species contribute significantly to particle growth during
606 these CLOUD chamber experiments. Water could play a role at higher relative humidities,
607 although quantum chemical calculations suggest that it plays only a minor role in NPF for the
608 system of sulfuric acid and dimethylamine (Olenius et al., 2017); this contrasts the sulfuric acid-
609 water system (see e.g. Zollner et al. 2012; Duplissy et al., 2016; Yu et al., 2017). In addition, it
610 is not exactly known how temperature influences the cluster evaporation rates (Hanson et al.,
611 2017). The evaporation rates from Jen et al. (2016a) were derived at temperatures close to 300

612 K; therefore the simulation of nucleation in the CLOUD chamber (278 K) using the Jen et al.
613 (2016a) rate parameters is likely to overestimate the effect of cluster evaporation.

614 It is not yet clear what exact base to acid ratio the particles have for a given diameter. The
615 clusters and small particles ($< \sim 2$ nm) seem to grow by maintaining a 1:1 ratio between base
616 and acid, which follows from measurements using mass spectrometers (Almeida et al., 2013;
617 Kürten et al., 2014; Bianchi et al., 2014). The larger particles could eventually reach a 2:1 ratio
618 between base and acid, especially at the DMA mixing ratios relevant for this study (Ahlm, et
619 al., 2016). However, even when a 2:1 ratio is assumed in the model (Fig. 2c) the expected size
620 distributions would not change significantly compared with the base-case scenario (1:1 ratio).
621 Therefore, it is not possible from our comparisons to find out if and at what diameter a transition
622 from 1:1 to 2:1 base to acid ratio takes place.

623 The question of why sulfuric acid-amine nucleation is rarely observed in the atmosphere is
624 still open. Jen et al. (2016a) reported that clusters that contain equal numbers of dimethylamine
625 and sulfuric acid molecules are ionized at reduced efficiencies than more acidic clusters with
626 the commonly used $\text{NO}_3^-(\text{HNO}_3)_{0.2}$ reagent ions. Still, Kürten et al. (2014) observed high
627 concentrations for large clusters containing acid and base at an average ratio of 1:1. A reduced
628 detection efficiency was also reported but the reduced sensitivity (in relation to the monomer)
629 was, e.g., only a factor of 3 for the trimer containing DMA. Using the model results from section
630 3.5 the expected trimer concentration at $5 \times 10^6 \text{ cm}^{-3}$ of sulfuric acid and 1 pptv ($2.5 \times 10^7 \text{ cm}^{-3}$)
631 of DMA should be $\sim 1 \times 10^5 \text{ cm}^{-3}$. Even when the detection efficiency for the trimer was reduced
632 by a factor of 3, such a concentration should still be well above the detection limit of a CI-API-
633 TOF. However, no sulfuric acid trimers could be detected in a field study where amines were
634 present at levels above 1 pptv ($2.5 \times 10^7 \text{ cm}^{-3}$, Kürten et al., 2016b). It is, therefore, possible that
635 any amines present were not suitable for nucleation. Therefore, application of methods capable
636 of amine speciation should be applied more widely in atmospheric measurements (Place et al.,
637 2017).

638 Several CLOUD papers reported particle formation rates for a diameter of 1.7 nm. Some of
639 these published formation rates were derived from direct measurements using particle counters
640 with cut-off diameters close to 1.7 nm (Riccobono et al., 2014; Duplissy et al., 2016), while
641 other reported NPF rates were derived from process models describing the nucleation process
642 in the CLOUD chamber (Kirkby et al., 2011; Kirkby et al., 2016). Therefore, no extrapolation
643 of the NPF rates from a larger threshold diameter was performed, which could have led to an
644 underestimation due to missing self-coagulation. Besides Almeida et al. (2013), the data set
645 reported by Dunne et al. (2016) and Kürten et al. (2016a) did make use of the NPF rate
646 extrapolation method from 3.2 to 1.7 nm without taking into account the effect of self-
647 coagulation. However, the reported formation rates are, in almost all cases, considerably slower
648 than those for the collision-controlled limit at a given sulfuric acid concentration since no
649 dimethylamine was present in the CLOUD chamber (Dunne et al., 2016; Kürten et al., 2016a).
650 The chemical system in these studies was the binary system, (H_2SO_4 and H_2O) and the ternary
651 system involving ammonia. The conditions only approached the collision-controlled limit at
652 the lowest temperature (210 K) when the highest ammonia mixing ratio of ~ 6 pptv (1.5×10^8
653 cm^{-3}) was investigated (Kürten et al., 2015b). However, even under these conditions, the
654 reported rates are only about a factor of 2 slower than the collision-controlled limit (Kürten et
655 al., 2016a). This is probably related to the low acid concentrations ($\leq 3 \times 10^6 \text{ cm}^{-3}$) in these

656 experiments, where the self-coagulation effect is not as strong as at higher acid concentration
657 (see Fig. 1) when wall loss and dilution lead to decreased cluster concentrations relative to the
658 monomer. This indicates that previously published CLOUD results, other than the Almeida et
659 al. (2013) data, are most likely not significantly affected.

660 McMurry and Li (2017) have recently investigated the effect of the wall loss and dilution
661 rate on new particle formation with their numerical model, which uses dimensionless
662 parameters. In order to allow for a comparison between McMurry and Li (2017) and the present
663 study, information on the dimensionless parameters W (describing wall loss) and M (describing
664 dilution) is provided (see McMurry and Li, 2017, for the exact definitions). These parameters
665 range from 0.04 to 0.7 (W) and 2×10^{-3} to 4×10^{-2} (M) for the experiments shown in this study
666 (Fig. 1). The monomer production rate (P_1) ranges from 7×10^3 to 2×10^6 $\text{cm}^{-3} \text{s}^{-1}$.

667
668
669

669 5. SUMMARY AND CONCLUSIONS

670
671 New particle formation rates from CLOUD chamber measurements for the sulfuric acid-
672 dimethylamine-water system were re-analyzed. It was found that the previously published rates
673 by Almeida et al. (2013) underestimate the NPF rates by up to a factor of ~ 50 at high acid
674 concentrations ($\sim 1 \times 10^7$ cm^{-3}). The reason for this underestimation is the effect of self-
675 coagulation that contributes efficiently to the loss of small particles in the size range relevant
676 for the data analysis (between 1.7 and 3.2 nm). The previously used method for extrapolating
677 the NPF rates from 3.2 nm to 1.7 nm did not include this effect and therefore the correction
678 factors were too small. Using an advanced reconstruction method that accounts for the effect
679 of self-coagulation yields much higher NPF rates (Kürten et al., 2015a). These corrected NPF
680 rates are in good agreement with rates calculated from an aerosol model assuming collision-
681 controlled nucleation and with measured NPF rates from SMPS data. Furthermore, the model
682 can reproduce the measured size distribution with good accuracy up to ~ 30 nm.

683 Extending the aerosol model by including evaporation rates for some clusters (see Jen et al.,
684 2016a) still yields good agreement between modeled and measured CLOUD NPF rates and
685 cluster concentrations. This indicates that the data for sulfuric acid-dimethylamine from the
686 flow tube study by Jen et al. (2016a) and from CLOUD (Kürten et al., 2014) are consistent for
687 the high base to acid ratio relevant for this study (dimethylamine to sulfuric acid monomer ratio
688 of ~ 100).

689 The above findings raise some further conclusions and questions. These are in part related
690 to the rare detection of sulfuric acid-amine nucleation in the atmosphere. Only one study has so
691 far reported sulfuric acid-amine nucleation (Zhao et al., 2011). The nucleation of sulfuric acid-
692 amines could occur, however, more often than currently thought.

693

694 – It is unclear to what extent previously published atmospheric NPF rates are affected by
695 incomplete J extrapolations. Some J measurements were made at diameters close to 3 nm
696 and extrapolated to a smaller size. If self-coagulation were important, the formation rates at
697 the small sizes could be significantly underestimated, and, therefore, in reality be much
698 closer to rates consistent with collision-controlled nucleation than previously thought. In
699 such a case, DMA (or other equally effective amines) could have been responsible for

700 nucleation as they are among the most potent nucleation precursors (in combination with
701 sulfuric acid). To avoid such ambiguities, the NPF rates should, in the future, be directly
702 measured at small diameters whenever possible.

703

704 – Better gas-phase amine (base) measurements are needed. Detection limits need to reach
705 mixing ratios even below 0.1 pptv ($2.5 \times 10^6 \text{ cm}^{-3}$); ideally the methods should also be
706 capable of speciating the amines (discriminate e.g. dimethylamine from ethylamine, which
707 have the same mass when measured by mass spectrometry but probably behave differently
708 in terms of their contribution to NPF). High time resolution (several minutes or better) for
709 the amine measurements during nucleation events is also important. This can show, whether
710 amines can be significantly depleted during NPF. As amines are not produced in the gas
711 phase (unlike sulfuric acid), their clustering with sulfuric acid monomers and small sulfuric
712 acid clusters/particles very likely can lead to a significant reduction in the amine mixing
713 ratios (Kürten et al., 2016b). This would indicate that new particle formation involving
714 amines in the atmosphere could be self-limiting, i.e., after an initial burst of particles, new
715 particle formation could be slowed down soon after when amine mixing ratios decrease.

716

717 – It is not clear why no clusters containing three or more sulfuric acid molecules are frequently
718 observed during atmospheric new particle formation when amines are expected to be
719 present. This could be due to incorrect assumptions about the amine concentrations, the
720 amine identities, or a reduced detection efficiency of chemical ionization mass
721 spectrometers (Jen et al., 2016a). The potential formation of complex multi-species clusters
722 (containing sulfuric acid, amines, ammonia and oxidized organics) in the atmosphere could
723 distribute the clusters over many different identities and therefore result in concentrations
724 too low to be detected by the current instrumentation for the individual species.

725

726 The overall contribution of amines to atmospheric nucleation can only be quantified after these
727 issues are understood. Besides further atmospheric measurements, controlled laboratory
728 measurements are necessary. Of special interest are the temperature dependent evaporation
729 rates of the relevant sulfuric-acid amine (and diamine) clusters.

730 **Appendix A:**

731

732 **Model that includes selected evaporation rates**

733

734 The kinetic model described in section 2.4 was expanded in a way that allows calculating the
 735 concentrations of the monomer, dimer, trimer and tetramer as a function of their dimethylamine
 736 content. Here, A_xB_y denotes the concentration of a cluster containing x sulfuric acid ($x = 1$ for
 737 the monomer) and y base ($y = 1$ for dimethylamine monomer) molecules; $x \geq y$ for all clusters,
 738 i.e., the number of bases is always smaller or equal to the number of acid molecules. When the
 739 total monomer concentration (N_1) is fixed, i.e., $A_1 = N_1 - A_1B_1$ at each time step, then the
 740 following equations result, i.e., for the A_1B_1 cluster

741

$$742 \frac{dA_1B_1}{dt} = K_{1,1} \cdot B_1 \cdot A_1 - (k_{1,w} + k_{dil} + k_{e,A_1B_1} + \sum_{j=1}^{N_{max}} K_{1,j} \cdot N_j) \cdot A_1B_1, \quad (A1)$$

743

744 for the two different identities of the sulfuric acid dimer

745

$$746 \frac{dA_2B_1}{dt} = (K_{1,1} \cdot A_1 \cdot A_1B_1 + k_{e,A_3B_1} \cdot A_3B_1) - (k_{w,2} + k_{dil} + K_{1,2} \cdot B_1 + \sum_{j=1}^N K_{j,2} \cdot N_j) \cdot$$

$$747 A_2B_1, \quad (A2)$$

748

$$749 \frac{dA_2B_2}{dt} = (0.5 \cdot K_{1,1} \cdot A_1B_1 \cdot A_1B_1 + K_{1,2} \cdot B_1 \cdot A_2B_1 + k_{e,A_3B_2} \cdot A_3B_2) - (k_{w,2} + k_{dil} +$$

$$750 \sum_{j=1}^N K_{j,2} \cdot N_j) \cdot A_2B_2, \quad (A3)$$

751

752 and for the three different identities of the sulfuric acid trimer

753

$$754 \frac{dA_3B_1}{dt} = (K_{1,2} \cdot A_1 \cdot A_2B_1) - (k_{w,3} + k_{dil} + k_{e,A_3B_1} + K_{1,3} \cdot B_1 + \sum_{j=1}^N K_{j,3} \cdot N_j - K_{1,3} \cdot A_1) \cdot$$

$$755 A_3B_1, \quad (A4)$$

756

$$757 \frac{dA_3B_2}{dt} = (K_{1,2} \cdot A_1B_1 \cdot A_2B_1 + K_{1,2} \cdot A_1 \cdot A_2B_2 + K_{1,3} \cdot B_1 \cdot A_3B_1) - (k_{w,3} + k_{dil} + k_{e,A_3B_2} +$$

$$758 K_{1,3} \cdot B_1 + \sum_{j=1}^N K_{j,3} \cdot N_j) \cdot A_3B_2, \quad (A5)$$

759

$$760 \frac{dA_3B_3}{dt} = (K_{1,2} \cdot A_1B_1 \cdot A_2B_2 + K_{1,3} \cdot B_1 \cdot A_3B_2) - (k_{w,3} + k_{dil} + \sum_{j=1}^N K_{j,3} \cdot N_j) \cdot A_3B_3. \quad (A6)$$

761

762 Since the formation of stable A_4B_1 clusters is not allowed (see Jen et al., 2016), the loss due to
 763 the A_1 and A_3B_1 collision is subtracted from the coagulation loss term in equation (A4).

764 Tetramers can be formed from trimers and dimers:

765

$$766 \frac{dN_4}{dt} = (K_{1,3} \cdot A_1B_1 \cdot A_3B_1 + K_{1,3} \cdot N_1 \cdot (A_3B_2 + A_3B_3) + 0.5 \cdot K_{2,2} \cdot N_2 \cdot N_2) - (k_{w,4} +$$

$$767 k_{dil} + \sum_{j=1}^N K_{j,4} \cdot N_j) \cdot N_4. \quad (A7)$$

768

769 Note that the formation of A_4B_1 (from A_3B_1) is not included in the formation rate for tetramers
 770 (see also further below). The concentrations of larger clusters and particles are calculated with

771 the same method as described in section 2.4. The cluster concentrations reported in section 3.5
772 refer to the number of acid molecules in the cluster, i.e., $N_1 = A_1 + A_1B_1$, $N_2 = A_2B_1 + A_2B_2$ and
773 $N_3 = A_3B_1 + A_3B_2 + A_3B_3$.

774 The evaporation rates considered are $k_{e,A_1B_1} = 0.1 \text{ s}^{-1}$, $k_{e,A_3B_1} = 1 \text{ s}^{-1}$, $k_{e,A_3B_2} = 1 \text{ s}^{-1}$ (Jen et al.,
775 2016a). Pure acid clusters are assumed to evaporate rapidly (at 278 K and higher) and are,
776 therefore, not considered (Hanson and Lovejoy, 2006). Jen et al. (2016a) suggested that the
777 formation of stable tetramers requires two base molecules. Therefore, this would indicate that
778 the evaporation rate k_{e,A_4B_1} is infinity (or very fast), which is also shown by Hanson et al. (2017).
779 However, the A_4B_1 formation (and its evaporation) is not explicitly treated in equations (A4)
780 and (A7).

781 In summary, three different evaporation rates were included in this model version (equations
782 (A1) to (A7)), i.e., $k_{e,A_1B_1} = 0.1 \text{ s}^{-1}$ (cluster A_1B_1), $k_{e,A_3B_1} = 1 \text{ s}^{-1}$ (cluster A_3B_1) and $k_{e,A_3B_2} = 1$
783 s^{-1} (cluster A_3B_2). All other evaporation rates were not explicitly included in the model, i.e.,
784 their rates were assumed to be zero (except for A_4B_1 , which is assumed to be infinity). Table 1
785 gives an overview of the different model configurations used to generate the model data in the
786 figures.

787

788

789 **Calculation of particle mobility diameters**

790

791 The mobility diameter of a cluster containing i sulfuric acid molecules (and i DMA molecules)
792 can be calculated according to

$$793 \quad d_{p,i} = \left(\frac{6 \cdot i \cdot M_w}{\pi \cdot N_A \cdot \rho} \right)^{1/3} + 0.3 \cdot 10^{-9} \text{ m}. \quad (\text{A8})$$

794

795 M_w is the molecular weight of the “monomer”, i.e., $0.143 \text{ kg mol}^{-1}$, ρ is the density of 1470 kg
796 m^{-3} (see section 2.4) and N_A is the Avogadro number, i.e., $6.022 \times 10^{23} \text{ mol}^{-1}$. The addition of
797 0.3 nm in equation (A8) is used to convert the geometric diameter (first term in equation (A8))
798 to a mobility diameter (Ku and Fernandez de la Mora, 2009).

799 **DATA AVAILABILITY**

800

801 Data used in this study can be obtained by sending an email to the corresponding author.

802

803

804 **ACKNOWLEDGEMENTS**

805

806 Funding from the German Federal Ministry of Education and Research (grant no. 01LK1222A)
807 and the Marie Curie Initial Training Network “CLOUD-TRAIN” (grant no. 316662) is
808 gratefully acknowledged. PHM's and CL's contributions to this work were supported by the US
809 Department of Energy’s Atmospheric System Research program, an Office of Science, Office
810 of Biological and Environmental Research, under grant number DE-SC0011780. RCF
811 acknowledges funding from the NSF Grants 1439551 and 1602086. MRP appreciates funding
812 from the Academy of Finland (project no. 299574). KL thanks the European Union’s Horizon
813 2020 research and innovation programme under the Marie Skłodowska-Curie grant agreement
814 no. 656994 (nano-CAVa).

815 **References**

- 816
- 817 Ahlm, L., Yli-Juuti, T., Schobesberger, S., Praplan, A. P., Kim, J., Tikkanen, O. -P., Lawler,
818 M. J., Smith, J. N., Tröstl, J., Acosta Navarro, J. C., Baltensperger, U., Bianchi, F., Donahue,
819 N. M., Duplissy, J., Franchin, A., Jokinen, T., Keskinen, H., Kirkby, J., Kürten, A., Laaksonen,
820 A., Lehtipalo, K., Petäjä, T., Riccobono, F., Rissanen, M. P., Rondo, L., Schallhart, S., Simon,
821 M., Winkler, P. M., Worsnop, D. R., Virtanen, A., and Riipinen, I.: Modeling the
822 thermodynamics and kinetics of sulfuric acid-dimethylamine-water nanoparticle growth in the
823 CLOUD chamber, *Aerosol Sci. Technol.*, 50, 1017–1032, doi:
824 10.1080/02786826.2016.1223268, 2016.
- 825
- 826 Almeida, J., Schobesberger, S., Kürten, A., Ortega, I. K., Kupiainen-Määttä, O., Praplan, A. P.,
827 Adamov, A., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Donahue,
828 N. M., Downard, A., Dunne, E. M., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Guida,
829 R., Hakala, J., Hansel, A., Heinritzi, M., Henschel, H., Jokinen, T., Junninen, H., Kajos, M.,
830 Kangasluoma, J., Keskinen, H., Kupc, A., Kurtén, T., Kvashin, A. N., Laaksonen, A., Lehtipalo,
831 K., Leiminger, M., Leppä, J., Loukonen, V., Makhmutov, V., Mathot, S., McGrath, M. J.,
832 Nieminen, T., Olenius, T., Onnela, A., Petäjä, T., Riccobono, F., Riipinen, I., Rissanen, M.,
833 Rondo, L., Ruuskanen, T., Santos, F. D., Sarnela, N., Schallhart, S., Schnitzhofer, R., Seinfeld,
834 J. H., Simon, M., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A., Tröstl, J., Tsagkogeorgas,
835 G., Vaattovaara, P., Viisanen, Y., Virtanen, A., Vrtala, A., Wagner, P. E., Weingartner, E.,
836 Wex, H., Williamson, C., Wimmer, D., Ye, P., Yli-Juuti, T., Carslaw, K. S., Kulmala, M.,
837 Curtius, J., Baltensperger, U., Worsnop, D. R., Vehkamäki, H., and Kirkby, J.: Molecular
838 understanding of sulphuric acid-amine particle nucleation in the atmosphere, *Nature*, 502, 359–
839 363, doi: 10.1038/nature12663, 2013.
- 840
- 841 Ball, S. M., Hanson, D. R., Eisele, F. L., and McMurry, P. H.: Laboratory studies of particle
842 nucleation: Initial results for H₂SO₄, H₂O, and NH₃ vapors, *J. Geophys. Res.-Atmos.*, 104, D19,
843 23709–23718, doi: 10.1029/1999JD900411, 1999.
- 844
- 845 Bergman, T., Laaksonen, A., Korhonen, H., Malila, J., Dunne, E. M., Mielonen, T., Lehtinen,
846 K. E. J., Kühn, T., Arola, A., and Kokkola, H.: Geographical and diurnal features of amine-
847 enhanced boundary layer nucleation, *J. Geophys. Res.-Atmos.*, 120, 9606–9624, doi:
848 10.1002/2015JD023181, 2015.
- 849
- 850 Berndt, T., Sipilä, M., Stratmann, F., Petäjä, T., Vanhanen, J., Mikkilä, J., Patokoski, J., Taipale,
851 R., Mauldin III, R. L., and Kulmala, M.: Enhancement of atmospheric H₂SO₄ / H₂O nucleation:
852 organic oxidation products versus amines, *Atmos. Chem. Phys.*, 14, 751–764, doi: 10.5194/acp-
853 14-751-2014, 2014.
- 854
- 855 Bianchi, F., Praplan, A. P., Sarnela, N., Dommen, J., Kürten, A., Ortega, I. K., Schobesberger,
856 S., Junninen, H., Simon, M., Tröstl, J., Jokinen, T., Sipilä, M., Adamov, A., Amorim, A.,
857 Almeida, J., Breitenlechner, M., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Hakala,
858 J., Hansel, A., Heinritzi, M., Kangasluoma, J., Keskinen, H., Kim, J., Kirkby, J., Laaksonen,

859 A., Lawler, M. J., Lehtipalo, K., Leiminger, M., Makhmutov, V., Mathot, S., Onnela, A., Petäjä,
860 T., Riccobono, F., Rissanen, M. P., Rondo, L., Tomé, A., Virtanen, A., Viisanen, Y.,
861 Williamson, C., Wimmer, D., Winkler, P. M., Ye, P., Curtius, J., Kulmala, M., Worsnop, D. R.,
862 Donahue, N. M., and Baltensperger, U.: Insight into acid-base nucleation experiments by
863 comparison of the chemical composition of positive, negative, and neutral clusters, *Environ.*
864 *Sci. Technol.*, 48, 13675–13684, doi: 10.1021/es502380b, 2014.

865

866 Chan, T. W., and Mozurkewich, M.: Measurement of the coagulation rate constant for sulfuric
867 acid particles as a function of particle size using tandem differential mobility analysis, *J.*
868 *Aerosol Sci.*, 32, 321–339, doi: 10.1016/S0021-8502(00)00081-1, 2001.

869

870 Chen, M., Titcombe, M., Jiang, J., Jen, C., Kuang, C., Fischer, M. L., Eisele, F. L., Siepmann,
871 J. I., Hanson, D. R., Zhao, J., and McMurry, P. H.: Acid–base chemical reaction model for
872 nucleation rates in the polluted atmospheric boundary layer, *P. Natl. Acad. Sci. USA*, 109,
873 18713–18718, doi: 10.1073/pnas.1210285109, 2012.

874

875 Clark, W. E., and Whitby, K. T.: Measurements of aerosols produced by the photochemical
876 oxidation of SO₂ in air, *J. Colloid Interface Sci.*, 51, 477–490, doi: 10.1016/0021-
877 9797(75)90144-7, 1975.

878

879 Crounse, J. D., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., and Wennberg, P. O.:
880 Autooxidation of organic compounds in the atmosphere, *J. Phys. Chem. Lett.*, 4, 3513–3520,
881 doi: 10.1021/jz4019207, 2013.

882

883 Crump, J. G., and Seinfeld, J. H.: Turbulent deposition and gravitational sedimentation of an
884 aerosol in a vessel of arbitrary shape, *J. Aerosol Sci.*, 12, 405–415, doi: 10.1016/0021-
885 8502(81)90036-7, 1981.

886

887 Dunne, E. M., Gordon, H., Kürten, A., Almeida, J., Duplissy, J., Williamson, C., Ortega, I. K.,
888 Pringle, K. J., Adamov, A., Baltensperger, U., Barmet, P., Benduhn, F., Bianchi, F.,
889 Breitenlechner, M., Clarke, A., Curtius, J., Dommen, J., Donahue, N. M., Ehrhart, S., Flagan,
890 R. C., Franchin, A., Guida, R., Hakala, J., Hansel, A., Heinritzi, M., Jokinen, T., Kangasluoma,
891 J., Kirkby, J., Kulmala, M., Kupc, A., Lawler, M. J., Lehtipalo, K., Makhmutov, V., Mann, G.,
892 Mathot, S., Merikanto, J., Miettinen, P., Nenes, A., Onnela, A., Rap, A., Reddington, C. L. S.,
893 Riccobono, F., Richards, N. A. D., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S.,
894 Sengupta, K., Simon, M., Sipilä, M., Smith, J. N., Stozkhov, Y., Tomé, A., Tröstl, J., Wagner,
895 P. E., Wimmer, D., Winkler, P. M., Worsnop, D. R., and Carslaw, K. S.: Global atmospheric
896 particle formation from CERN CLOUD measurements, *Science*, 354, 1119–1124, doi:
897 10.1126/science.aaf2649, 2016.

898

899 Duplissy, J., Merikanto, J., Franchin, A., Tsagkogeorgas, G., Kangasluoma, J., Wimmer, D.,
900 Vuollekoski, H., Schobesberger, S., Lehtipalo, K., Flagan, R. C., Brus, D., Donahue, N. M.,
901 Vehkämäki, H., Almeida, J., Amorim, A., Barmet, P., Bianchi, F., Breitenlechner, M., Dunne,
902 E. M., Guida, R., Henschel, H., Junninen, H., Kirkby, J., Kürten, A., Kupc, A., Määttänen, A.,

903 Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Praplan, A. P., Riccobono, F., Rondo,
904 L., Steiner, G., Tome, A., Walther, H., Baltensperger, U., Carslaw, K. S., Dommen, J., Hansel,
905 A., Petäjä, T., Sipilä, M., Stratmann, F., Vrtala, A., Wagner, P. E., Worsnop, D. R., Curtius, J.,
906 and Kulmala, M.: Effect of ions on sulfuric acid-water binary particle formation II:
907 Experimental data and comparison with QC-normalized classical nucleation theory, *J.*
908 *Geophys. Res.-Atmos.*, 121, 1752–1775, doi: 10.1002/2015JD023539, 2016.

909

910 Ehrhart, S. and Curtius, J.: Influence of aerosol lifetime on the interpretation of nucleation
911 experiments with respect to the first nucleation theorem, *Atmos. Chem. Phys.*, 13, 11465–
912 11471, doi: 10.5194/acp-13-11465-2013, 2013.

913

914 Ehrhart, S., Ickes, L., Almeida, J., Amorim, A., Barmet, P., Bianchi, F., Dommen, J., Dunne,
915 E. M., Duplissy, J., Franchin, A., Kangasluoma, J., Kirkby, J., Kürten, A., Kupc, A., Lehtipalo,
916 K., Nieminen, T., Riccobono, F., Rondo, L., Schobesberger, S., Steiner, G., Tomé, A., Wimmer,
917 D., Baltensperger, U., Wagner, P. E., and Curtius, J.: Comparison of the SAWNUC model with
918 CLOUD measurements of sulphuric acid-water nucleation, *J. Geophys. Res.-Atmos.*, 121,
919 12401–12414, doi: 10.1002/2015JD023723, 2016.

920

921 Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach,
922 F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T.,
923 Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen, L. B.,
924 Jørgensen, S., Kjaergaard, H. G., Canagaratna, M., Dal Maso, M., Berndt, T., Petäjä, T.,
925 Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A
926 large source of low-volatility secondary organic aerosol, *Nature*, 506, 476–479, doi:
927 10.1038/nature13032, 2014.

928

929 Fiedler, V., Dal Maso, M., Boy, M., Aufmhoff, H., Hoffmann, J., Schuck, T., Birmili, W.,
930 Hanke, M., Uecker, J., Arnold, F., and Kulmala, M.: The contribution of sulphuric acid to
931 atmospheric particle formation and growth: a comparison between boundary layers in Northern
932 and Central Europe, *Atmos. Chem. Phys.*, 5, 1773–1785, doi: 10.5194/acp-5-1773-2005, 2005.

933

934 Freshour, N. A., Carlson, K. K., Melka, Y. A., Hinz, S., Panta, B., and Hanson, D. R.: Amine
935 permeation sources characterized with acid neutralization and sensitivities of an amine mass
936 spectrometer, *Atmos. Meas. Tech.*, 7, 3611–3621, doi: 10.5194/amt-7-3611-2014, 2014.

937

938 Glasoe, W. A., Volz, K., Panta, B., Freshour, N., Bachman, R., Hanson, D. R., McMurry, P.
939 H., and Jen, C.: Sulfuric acid nucleation: An experimental study of the effect of seven bases, *J.*
940 *Geophys. Res.-Atmos.*, 120, 1933–1950, doi: 10.1002/2014JD022730, 2015.

941

942 Gordon, H., Sengupta, K., Rap, A., Duplissy, J., Frege, C., Williamson, C., Heinritzi, M.,
943 Simon, M., Yan, C., Almeida, J., Tröstl, J., Nieminen, T., Ortega, I. K., Wagner, R., Dunne, E.
944 M., Adamov, A., Amorim, A., Bernhammer, A. K., Bianchi, F., Breitenlechner, M., Brilke, S.,
945 Chen, X., Craven, J. S., Dias, A., Ehrhart, S., Fischer, L., Flagan, R. C., Franchin, A., Fuchs,
946 C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim, J.,

947 Kirkby, J., Krapf, M., Kürten, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S.,
948 Molteni, U., Monks, S. A., Onnela, A., Peräkylä, O., Piel, F., Petäjä, T., Praplan, A. P., Pringle,
949 K. J., Richards, N. A. D., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Scott, C.
950 E., Seinfeld, J. H., Sharma, S., Sipilä, M., Steiner, G., Stozhkov, Y., Stratmann, F., Tomé, A.,
951 Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D.,
952 Winkler, P. M., Ye, P., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R.,
953 Baltensperger, U., Kulmala, M., Curtius, J., and Carslaw, K. S.: Reduced anthropogenic aerosol
954 radiative forcing caused by biogenic new particle formation, *Proc. Natl. Acad. Sci. USA*, 113,
955 12053–12058, doi: 10.1073/pnas.1602360113, 2016.

956

957 Hamaker, H. C.: The London–van der Waals attraction between spherical particles, *Physica*, 4,
958 1058–1072, doi: 10.1016/S0031-8914(37)80203-7, 1937.

959

960 Hanson, D. R., and Eisele, F.: Diffusion of H₂SO₄ in humidified nitrogen: Hydrated H₂SO₄, *J.*
961 *Phys. Chem. A*, 104, 1715–1719, doi: 10.1021/jp993622j, 2000.

962

963 Hanson, D. R., and Lovejoy, E. R.: Measurement of the thermodynamics of the hydrated dimer
964 and trimer of sulfuric acid, *J. Phys. Chem. A*, 110, 9525–9528, doi: 10.1021/jp062844w, 2006.

965

966 Hanson, D. R., Bier, I., Panta, B., Jen, C. N., and McMurry, P. H.: Computational Fluid
967 Dynamics Studies of a Flow Reactor: Free Energies of Clusters of Sulfuric Acid with NH₃ or
968 Dimethyl Amine, *J. Phys. Chem. A*, 121, 3976–3990, doi: 10.1021/acs.jpca.7b00252, 2017.

969

970 Hinds, W. C.: Aerosol technology: Properties, behavior, and measurement of airborne particles,
971 second edition, John Wiley & Sons, Inc., 150–153, 1999.

972

973 Hirsikko, A., Laakso, L., Hörrak, U., Aalto, P. P., Kerminen, V.-M., and Kulmala, M.: Annual
974 and size dependent variation of growth rates and ion concentrations in boreal forest, *Boreal*
975 *Environ. Res.*, 10, 357–369, 2005.

976

977 Jen, C., McMurry, P. H., and Hanson, D. R.: Stabilization of sulfuric acid dimers by ammonia,
978 methylamine, dimethylamine, and trimethylamine, *J. Geophys. Res.-Atmos.*, 119, 7502–7514,
979 doi: 10.1002/2014JD021592, 2014.

980

981 Jen, C. N., Zhao, J., McMurry, P. H., and Hanson, D. R.: Chemical ionization of clusters formed
982 from sulfuric acid and dimethylamine or diamines, *Atmos. Chem. Phys.*, 16, 12513–12529, doi:
983 10.5194/acp-16-12513-2016, 2016a.

984

985 Jen, C. N., Bachman, R., Zhao, J., McMurry, P. H., and Hanson, D. R.: Diamine-sulfuric acid
986 reactions are a potent source of new particle formation, *Geophys. Res. Lett.*, 43, 867–873, doi:
987 10.1002/2015GL066958, 2016b.

988

989 Jokinen, T., Sipilä, M., Junninen, H., Ehn, M., Lönn, G., Hakala, J., Petäjä, T., Mauldin III, R.
990 L., Kulmala, M., and Worsnop, D. R.: Atmospheric sulphuric acid and neutral cluster

991 measurements using CI-API-TOF, *Atmos. Chem. Phys.*, 12, 4117–4125, doi: 10.5194/acp-12-
992 4117-2012, 2012.

993

994 Jokinen, T., Berndt, T., Makkonen, R., Kerminen, V.-M., Junninen, H., Paasonen, P.,
995 Stratmann, F., Herrmann, H., Guenther, A. B., Worsnop, D. R., Kulmala, M., Ehn, M., and
996 Sipilä, M.: Production of extremely low volatile organic compounds from biogenic emissions:
997 Measured yields and atmospheric implications, *P. Natl. Acad. Sci. USA*, 112, 7123–7128, doi:
998 10.1073/pnas.1423977112, 2015.

999

1000 Karlsson, M. N. A., and Martinsson, B. G.: Methods to measure and predict the transfer function
1001 size dependence of individual DMAs, *J. Aerosol Sci.*, 34, 603–625, doi: 10.1016/S0021-
1002 8502(03)00020-X, 2003.

1003

1004 Kerminen, V.-M., and Kulmala, M.: Analytical formulae connecting the “real” and the
1005 “apparent” nucleation rate and the nuclei number concentration for atmospheric nucleation
1006 events, *J. Aerosol Sci.*, 33, 609–622, doi: 10.1016/S0021-8502(01)00194-X, 2002.

1007

1008 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné,
1009 S., Ickes, L., Kürten, A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S.,
1010 Tsagkogeorgas, G., Wimmer, D., Amorim, A., Bianchi, F., Breitenlechner, M., David, A.,
1011 Dommen, J., Downard, A., Ehn, M., Flagan, R.C., Haider, S., Hansel, A., Hauser, D., Jud, W.,
1012 Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R.,
1013 Makhmutov, V., Mathot, S., Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A.,
1014 Pereira, P., Petäjä, T., Schnitzhofer, R., Seinfeld, J. H., Sipilä, M., Stozhkov, Y., Stratmann, F.,
1015 Tomé, A., Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E.,
1016 Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U., and Kulmala, M.:
1017 Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation,
1018 *Nature*, 476, 429–435, doi: 10.1038/nature10343, 2011.

1019

1020 Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M.,
1021 Simon, M., Yan, C., Almeida, J., Tröstl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov,
1022 A., Amorim, A., Bernhammer, A.-K., Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X.,
1023 Craven, J., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Hakala, J.,
1024 Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim, J., Krapf, M., Kürten, A.,
1025 Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Peräkylä,
1026 O., Piel, F., Petäjä, T., Praplan, A. P., Pringle, K., Rap, A., Richards, N. A. D., Riipinen, I.,
1027 Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Scott, C. E., Seinfeld, J. H., Sipilä,
1028 M., Steiner, G., Stozhkov, Y., Stratmann, F., Tomé, A., Virtanen, A., Vogel, A. L., Wagner, A.,
1029 Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Zhang, X., Hansel, A.,
1030 Dommen, J., Donahue, N. M., Worsnop, D. R., Baltensperger, U., Kulmala, M., Carslaw, K.
1031 S., and Curtius, J.: Ion-induced nucleation of pure biogenic particles, *Nature*, 533, 521–526,
1032 doi: 10.1038/nature17953, 2016.

1033

1034 Ku, B. K., and Fernandez de la Mora, J.: Relation between electrical mobility, mass, and size
1035 for nanodrops 1–6.5 nm in diameter in air, *Aerosol Sci. Technol.*, 43, 241–249, doi:
1036 10.1080/02786820802590510, 2009.
1037
1038 Kuang, C., McMurry, P. H., McCormick, A. V., and Eisele, F. L.: Dependence of nucleation
1039 rates on sulfuric acid vapor concentration in diverse atmospheric locations, *J. Geophys. Res.-*
1040 *Atmos.*, 113, D10209, doi: 10.1029/2007JD009253, 2008.
1041
1042 Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V.-M., Birmili,
1043 W., and McMurry, V.-M.: Formation and growth rates of ultrafine atmospheric particles: a
1044 review of observations, *J. Aerosol Sci.*, 35, 143–176, doi: 10.1016/j.jaerosci.2003.10.003, 2004.
1045
1046 Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T.,
1047 Petäjä, T., Sipilä, M., Schobesberger, S., Rantala, P., Franchin, A., Jokinen, T., Järvinen, E.,
1048 Äijälä, M., Kangasluoma, J., Hakala, J., Aalto, P. P., Paasonen, P., Mikkilä, J., Vanhanen, J.,
1049 Aalto, J., Hakola, H., Makkonen, U., Ruuskanen, T., Mauldin III, R. L., Duplissy, J.,
1050 Vehkamäki, H., Bäck, J., Kortelainen, A., Riipinen, I., Kurtén, T., Johnston, M. V., Smith, J.
1051 N., Ehn, M., Mentel, T. F., Lehtinen, K. E. J., Laaksonen, A., Kerminen, V.-M., and Worsnop,
1052 D. R.: Direct observations of atmospheric aerosol nucleation, *Science*, 339, 943–946, doi:
1053 10.1126/science.1227385, 2013.
1054
1055 Kupc, A., Amorim, A., Curtius, J., Danielczok, A., Duplissy, J., Ehrhart, S., Walther, H., Ickes,
1056 L., Kirkby, J., Kürten, A., Lima, J. M., Mathot, S., Minginette, P., Onnela, A., Rondo, L., and
1057 Wagner, P. E.: A fibre-optic UV system for H₂SO₄ production in aerosol chambers causing
1058 minimal thermal effects, *J. Aerosol Sci.*, 42, 532–543, doi: 10.1016/j.jaerosci.2011.05.001,
1059 2011.
1060
1061 Kürten, A., Rondo, L., Ehrhart, S., and Curtius, J.: Performance of a corona ion source for
1062 measurement of sulfuric acid by chemical ionization mass spectrometry, *Atmos. Meas. Tech.*,
1063 4, 437–443, doi: 10.5194/amt-4-437-2011, 2011.
1064
1065 Kürten, A., Jokinen, T., Simon, M., Sipilä, M., Sarnela, N., Junninen, H., Adamov, A., Almeida,
1066 J., Amorim, A., Bianchi, F., Breitenlechner, M., Dommen, J., Donahue, N. M., Duplissy, J.,
1067 Ehrhart, S., Flagan, R. C., Franchin, A., Hakala, J., Hansel, A., Heinritzi, M., Hutterli, M.,
1068 Kangasluoma, J., Kirkby, J., Laaksonen, A., Lehtipalo, K., Leiminger, M., Makhmutov, V.,
1069 Mathot, S., Onnela, A., Petäjä, T., Praplan, A. P., Riccobono, F., Rissanen, M. P., Rondo, L.,
1070 Schobesberger, S., Seinfeld, J. H., Steiner, G., Tomé, A., Tröstl, J., Winkler, P. M., Williamson,
1071 C., Wimmer, D., Ye, P., Baltensperger, U., Carslaw, K. S., Kulmala, M., Worsnop, D. R., and
1072 Curtius, J.: Neutral molecular cluster formation of sulfuric acid-dimethylamine observed in
1073 real-time under atmospheric conditions, *P. Natl. Acad. Sci. USA*, 111, 15019–15024, doi:
1074 10.1073/pnas.1404853111, 2014.
1075

1076 Kürten, A., Williamson, C., Almeida, J., Kirkby, J., and Curtius, J.: On the derivation of particle
1077 nucleation rates from experimental formation rates, *Atmos. Chem. Phys.*, 15, 4063–4075, doi:
1078 10.5194/acp-15-4063-2015, 2015a.
1079

1080 Kürten, A., Münch, S., Rondo, L., Bianchi, F., Duplissy, J., Jokinen, T., Junninen, H., Sarnela,
1081 N. Schobesberger, S., Simon, M., Sipilä, M., Almeida, J., Amorim, A., Dommen, J., Donahue,
1082 N. M., Dunne, M., Flagan, R. C., Franchin, A., Kirkby, J., Kupc, A., Makhmutov, V., Petäjä,
1083 T. Praplan, A. P., Riccobono, F., Steiner, G., Tomé, A., Tsagkogeorgas, G., Wagner, P. E.,
1084 Wimmer, D., Baltensperger, U., Kulmala, M., Worsnop, D. R., and Curtius, J.:
1085 Thermodynamics of the formation of sulfuric acid dimers in the binary (H₂SO₄-H₂O) and
1086 ternary (H₂SO₄-H₂O-NH₃) system, *Atmos. Chem. Phys.*, 15, 10701–10721, doi: 10.5194/acp-
1087 15-10701-2015, 2015b.
1088

1089 Kürten, A., Bianchi, F., Almeida, J., Kupiainen-Määttä, O., Dunne, E. M., Duplissy, J.,
1090 Williamson, C., Barmet, P., Breitenlechner, M., Dommen, J., Donahue, N. M., Flagan, R. C.,
1091 Franchin, A., Gordon, H., Hakala, J., Hansel, A., Heinritzi, M., Ickes, L., Jokinen, T.,
1092 Kangasluoma, J., Kim, J., Kirkby, J., Kupc, A., Lehtipalo, K., Leiminger, M., Makhmutov, V.,
1093 Onnela, A., Ortega, I. K., Petäjä, T., Praplan, A. P., Riccobono, F., Rissanen, M. P., Rondo, L.,
1094 Schnitzhofer, R., Schobesberger, S., Smith, J. N., Steiner, G., Stozhkov, Y., Tomé, A., Tröstl,
1095 J., Tsagkogeorgas, G., Wagner, P. E., Wimmer, D., Ye, P., Baltensperger, U., Carslaw, K.,
1096 Kulmala, M., and Curtius, J.: Experimental particle formation rates spanning tropospheric
1097 sulfuric acid and ammonia abundances, ion production rates and temperatures, *J. Geophys.*
1098 *Res.-Atmos.*, 121, 12377–12400, doi: 10.1002/2015JD023908, 2016a.
1099

1100 Kürten, A., Bergen, A., Heinritzi, M., Leiminger, M., Lorenz, V., Piel, F., Simon, M., Sitals,
1101 R., Wagner, A. C., and Curtius, J.: Observation of new particle formation and measurement of
1102 sulfuric acid, ammonia, amines and highly oxidized organic molecules at a rural site in central
1103 Germany, *Atmos. Chem. Phys.*, 16, 12793–12813, doi: 10.5194/acp-16-12793-2016, 2016b.
1104

1105 Kurtén, T., Loukonen, V., Vehkamäki, H., and Kulmala, M.: Amines are likely to enhance
1106 neutral and ion-induced sulfuric acid-water nucleation in the atmosphere more effectively than
1107 ammonia, *Atmos. Chem. Phys.*, 8, 4095–4103, doi: 10.5194/acp-8-4095-2008, 2008.
1108

1109 Lee, S.-H., Reeves, J. M., Wilson, J. C., Hunton, D. E., Viggiano, A. A., Miller, T. M.,
1110 Ballenthin, J. O., and Lait, L. R.: Particle formation by ion nucleation in the upper troposphere
1111 and lower stratosphere, *Science*, 301, 1886–1889, doi: 10.1126/science.1087236, 2003.
1112

1113 Lehtipalo, K., Rondo, L., Kontkanen, J., Schobesberger, S., Jokinen, T., Sarnela, N., Kürten,
1114 A., Ehrhart, S., Franchin, A., Nieminen, T., Riccobono, F., Sipilä, M., Yli-Juuti, T., Duplissy,
1115 J., Adamov, A., Ahlm, L., Almeida, J., Amorim, A., Bianchi, F., Breitenlechner, M., Dommen,
1116 J., Downard, A. J., Dunne, E. M., Flagan, R. C., Guida, R., Hakala, J., Hansel, A., Jud, W.,
1117 Kangasluoma, J., Kerminen, V.-M., Keskinen, H., Kim, J., Kirkby, J., Kupc, A., Kupiainen-
1118 Määttä, O., Laaksonen, A., Lawler, M. J., Leiminger, M., Mathot, S., Olenius, T., Ortega, I. K.,
1119 Onnela, A., Petäjä, T., Praplan, A., Rissanen, M. P., Ruuskanen, T., Santos, F. D., Schallhart,

1120 S., Schnitzhofer, R., Simon, M., Smith, J. N., Tröstl, J., Tsagkogeorgas, G., Tomé, A.,
1121 Vaattovaara, P., Vehkamäki, H., Vrtala, A. E., Wagner, P. E., Williamson, C., Wimmer, D.,
1122 Winkler, P. M., Virtanen, A., Donahue, N. M., Carslaw, K. S., Baltensperger, U., Riipinen, I.,
1123 Curtius, J., Worsnop, D. R., and Kulmala, M.: The effect of acid–base clustering and ions on
1124 the growth of atmospheric nano-particles, *Nature Commun.*, 7, 11594, doi:
1125 10.1038/ncomms11594, 2016.
1126
1127 Lovejoy, E. R., Curtius, J., and Froyd, K. D.: Atmospheric ion-induced nucleation of sulfuric
1128 acid and water, *J. Geophys. Res.-Atmos.*, 109, D08204, doi: 10.1029/2003JD004460, 2004.
1129
1130 McGrath, M. J., Olenius, T., Ortega, I. K., Loukonen, V., Paasonen, P., Kurtén, T., Kulmala,
1131 M., and Vehkamäki, H.: Atmospheric Cluster Dynamics Code: a flexible method for solution
1132 of the birth-death equations, *Atmos. Chem. Phys.*, 12, 2345–2355, doi: 10.5194/acp-12-2345-
1133 2012, 2012.
1134
1135 McMurry, P. H.: Photochemical Aerosol Formation from SO₂: A theoretical analysis of smog
1136 chamber data, *J. Colloid Interf. Sci.*, 78, 513–527, doi: 10.1016/0021-9797(80)90589-5, 1980.
1137
1138 McMurry, P. H., and Li, C.: The dynamic behavior of nucleating aerosols in constant reaction
1139 rate systems: Dimensional analysis and generic numerical solutions, *Aerosol Sci. Technol.*, 51,
1140 1057–1070, doi: 10.1080/02786826.2017.1331292, 2017.
1141
1142 Nadykto, A. B., Yu, F., Jakovleva, M. V., Herb, J., and Xu, Y.: Amines in the Earth’s
1143 atmosphere: A density functional theory study of the thermochemistry of pre-nucleation
1144 clusters, *Entropy*, 13, 554–569; doi: 10.3390/e13020554, 2011.
1145
1146 Olenius, T., Halonen, R., Kurtén, T., Henschel, H., Kupiainen-Määttä, O., Ortega, I. K., Jen, C.
1147 N., Vehkamäki, H., and Riipinen, I.: New particle formation from sulfuric acid and amines:
1148 Comparison of monomethylamine, dimethylamine, and trimethylamine, *J. Geophys. Res.*
1149 *Atmos.*, 122, 7103–7118, doi: 10.1002/2017JD026501, 2017.
1150
1151 Ortega, I. K., Kupiainen, O., Kurtén, T., Olenius, T., Wilkman, O., McGrath, M. J., Loukonen,
1152 V., and Vehkamäki, H.: From quantum chemical formation free energies to evaporation rates,
1153 *Atmos. Chem. Phys.*, 12, 225–235, doi: 10.5194/acp-12-225-2012, 2012.
1154
1155 Park, S. H., Kim, H. O., Han, Y. T., Kwon, S. B., and Lee, K. W.: Wall loss rate of polydispersed
1156 aerosols, *Aerosol Sci. Technol.*, 35, 710–717, doi: 10.1080/02786820152546752, 2001.
1157
1158 Place, B. K., Quilty, A. T., Di Lorenzo, R. A., Ziegler, S. E., and VandenBoer, T. C.:
1159 Quantitation of 11 alkylamines in atmospheric samples: separating structural isomers by ion
1160 chromatography, *Atmos. Meas. Tech.*, 10, 1061–1078, doi: 10.5194/amt-10-1061-2017, 2017.
1161

1162 Praplan, A. P., Bianchi, F., Dommen, J., and Baltensperger, U.: Dimethylamine and ammonia
1163 measurements with ion chromatography during the CLOUD4 campaign, *Atmos. Meas. Tech.*,
1164 5, 2161–2167, doi: 10.5194/amt-5-2161-2012, 2012.
1165

1166 Rao, N. P., and McMurry, P. H.: Nucleation and Growth of Aerosol in Chemically Reacting
1167 Systems: A Theoretical Study of the Near-Collision-Controlled Regime, *Aerosol Sci. Technol.*,
1168 11, 120–132, doi: 10.1080/02786828908959305, 1989.
1169

1170 Qiu, C., and Zhang, R.: Physiochemical properties of alkylammonium sulfates: hygroscopicity,
1171 thermostability, and density, *Environ. Sci. Technol.*, 46, 4474–4480, doi: 10.1021/es3004377,
1172 2012.
1173

1174 Riccobono, F., Schobesberger, S., Scott, C. E., Dommen, J., Ortega, I. K., Rondo, L., Almeida,
1175 J., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Downard, A., Dunne, E. M.,
1176 Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Hansel, A., Junninen, H., Kajos, M.,
1177 Keskinen, H., Kupc, A., Kürten, A., Kvashin, A. N., Laaksonen, A., Lehtipalo, K., Makhmutov,
1178 V., Mathot, S., Nieminen, T., Onnela, A., Petäjä, T., Praplan, A. P., Santos, F. D., Schallhart,
1179 S., Seinfeld, J. H., Sipilä, M., Spracklen, D. V., Stozhkov, Y., Stratmann, F., Tomé, A.,
1180 Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y., Vrtala, A., Wagner, P. E., Weingartner, E.,
1181 Wex, H., Wimmer, D., Carslaw, K. S., Curtius, J., Donahue, N. M., Kirkby, J., Kulmala, M.,
1182 Worsnop, D. R., and Baltensperger, U.: Oxidation products of biogenic emissions contribute to
1183 nucleation of atmospheric particles, *Science*, 344, 717–721, doi: 10.1126/science.1243527,
1184 2014.
1185

1186 Simon, M., Heinritzi, M., Herzog, S., Leiminger, M., Bianchi, F., Praplan, A., Dommen, J.,
1187 Curtius, J., and Kürten, A.: Detection of dimethylamine in the low pptv range using nitrate
1188 chemical ionization atmospheric pressure interface time-of-flight (CI-API-TOF) mass
1189 spectrometry, *Atmos. Meas. Tech.*, 9, 2135–2145, doi: 10.5194/amt-9-2135-2016, 2016.
1190

1191 Voigtländer, J., Duplissy, J., Rondo, L., Kürten, A., and Stratmann, F.: Numerical simulations
1192 of mixing conditions and aerosol dynamics in the CERN CLOUD chamber, *Atmos. Chem.*
1193 *Phys.*, 12, 2205–2214, doi: 10.5194/acp-12-2205-2012, 2012.
1194

1195 Wang, S. C., and Flagan, R. C.: Scanning electrical mobility spectrometer, *Aerosol Sci.*
1196 *Technol.*, 13, 230–240, doi: 10.1080/02786829008959441, 1990.
1197

1198 Weber, R. J., Marti, J., McMurry, P. H., Eisele, F. L., Tanner, D. J., and Jefferson, A.: Measured
1199 atmospheric new particle formation rates: implications for nucleation mechanisms, *Chem. Eng.*
1200 *Comm.*, 151, 53–64, doi: 10.1080/00986449608936541, 1996.
1201

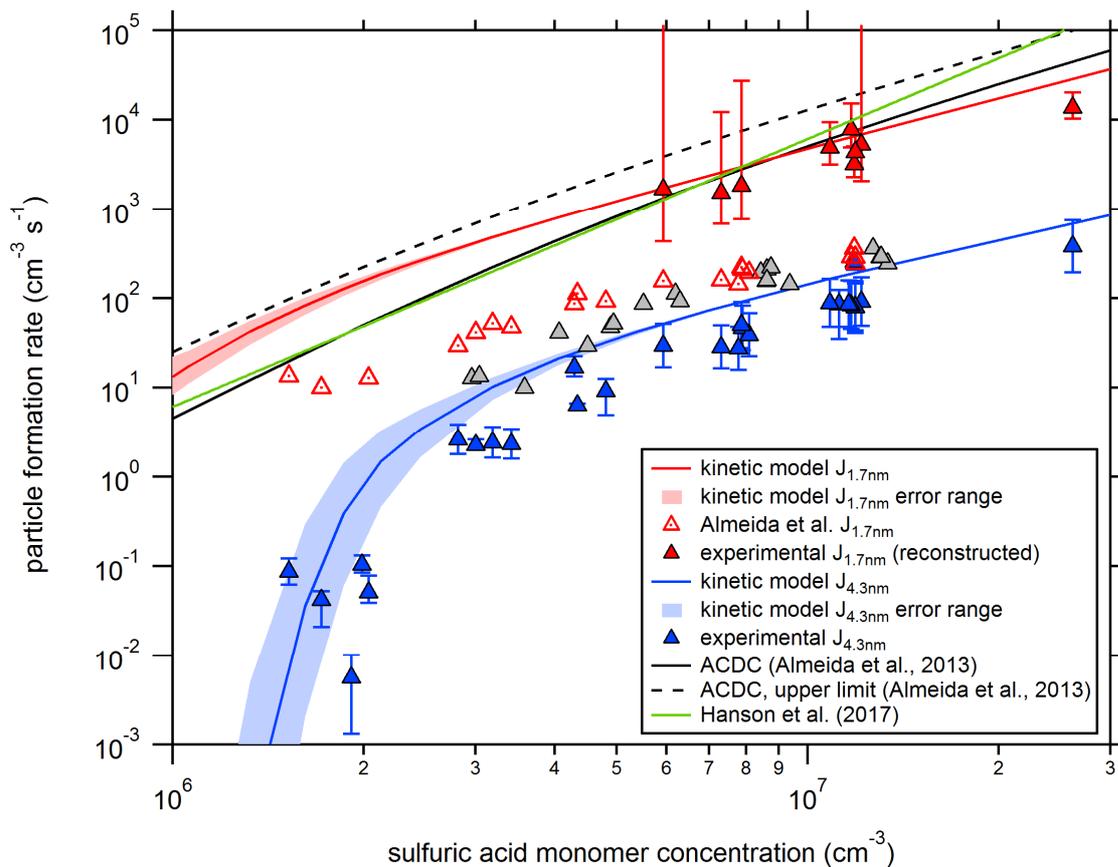
1202 Weber, R. J., Marti, J. J., McMurry, P. H., Eisele, F. L., Tanner, D. J., and Jefferson, A.:
1203 Measurements of new particle formation and ultrafine particle growth rates at a clean
1204 continental site, *J. Geophys. Res.-Atmos.*, 102, D4, 4375–4385, doi: 10.1029/96JD03656, 1997.
1205

1206 Weber, R. J., McMurry, P. H., Mauldin, L., Tanner, D. J., Eisele, F. L., Brechtel, F. J.,
1207 Kreidenweis, S. M., Kok, G. L., Schillawski, R. D., and Baumgardner, D.: A study of new
1208 particle formation and growth involving biogenic and trace gas species measured during ACE
1209 1, *J. Geophys. Res.-Atmos.*, 103, D13, 16385–16396, doi: 10.1029/97JD02465, 1998.
1210
1211 Wiedensohler, A., and Fissan, H. J.: Aerosol charging in high purity gases, *J. Aerosol Sci.*, 19,
1212 867–870, doi: 10.1016/0021-8502(88)90054-7, 1988.
1213
1214 Yao, L., Wang, M.-Y., Wang, X.-K., Liu, Y.-J., Chen, H.-F., Zheng, J., Nie, W., Ding, A.-J.,
1215 Geng, F.-H., Wang, D.-F., Chen, J.-M., Worsnop, D. R., and Wang, L.: Detection of
1216 atmospheric gaseous amines and amides by a high-resolution time-of-flight chemical ionization
1217 mass spectrometer with protonated ethanol reagent ions, *Atmos. Chem. Phys.*, 16, 14527–
1218 14543, doi: 10.5194/acp-16-14527-2016, 2016.
1219
1220 You, Y., Kanawade, V. P., de Gouw, J. A., Guenther, A. B., Madronich, S., Sierra-Hernández,
1221 M. R., Lawler, M., Smith, J. N., Takahama, S., Ruggeri, G., Koss, A., Olson, K., Baumann, K.,
1222 Weber, R. J., Nenes, A., Guo, H., Edgerton, E. S., Porcelli, L., Brune, W. H., Goldstein, A. H.,
1223 and Lee, S.-H.: Atmospheric amines and ammonia measured with a chemical ionization mass
1224 spectrometer (CIMS), *Atmos. Chem. Phys.*, 14, 12181–12194, doi: 10.5194/acp-14-12181-
1225 2014, 2014.
1226
1227 Yu, H., and Lee, S.-H.: Chemical ionisation mass spectrometry for the measurement of
1228 atmospheric amines, *Environ. Chem.*, 9, 190–201, doi: 10.1071/EN12020, 2012.
1229
1230 Yu, H., Dai, L., Zhao, Y., Kanawade, V. P., Tripathi, S. N., Ge, X., Chen, M., and Lee, S. N.:
1231 Laboratory observations of temperature and humidity dependencies of nucleation and growth
1232 rates of sub-3 nm particles, *J. Geophys. Res. Atmos.*, 122, 1919–1929, doi:
1233 10.1002/2016JD025619, 2017.
1234
1235 Zhao, J., Smith, J. N., Eisele, F. L., Chen, M., Kuang, C., and McMurry, P. H.: Observation of
1236 neutral sulfuric acid-amine containing clusters in laboratory and ambient measurements, *Atmos.*
1237 *Chem. Phys.*, 11, 10823–10836, doi: 10.5194/acp-11-10823-2011, 2011.
1238
1239 Zollner, J. H., Glasoe, W. A., Panta, B., Carlson, K. K., McMurry, P. H., and Hanson, D. R.:
1240 Sulfuric acid nucleation: power dependencies, variation with relative humidity, and effect of
1241 bases, *Atmos. Chem. Phys.*, 12, 4399–4411, doi: 10.5194/acp-12-4399-2012, 2012.

1242 **Table 1.** Overview of the two different model versions used to generate the data in the figures.
 1243

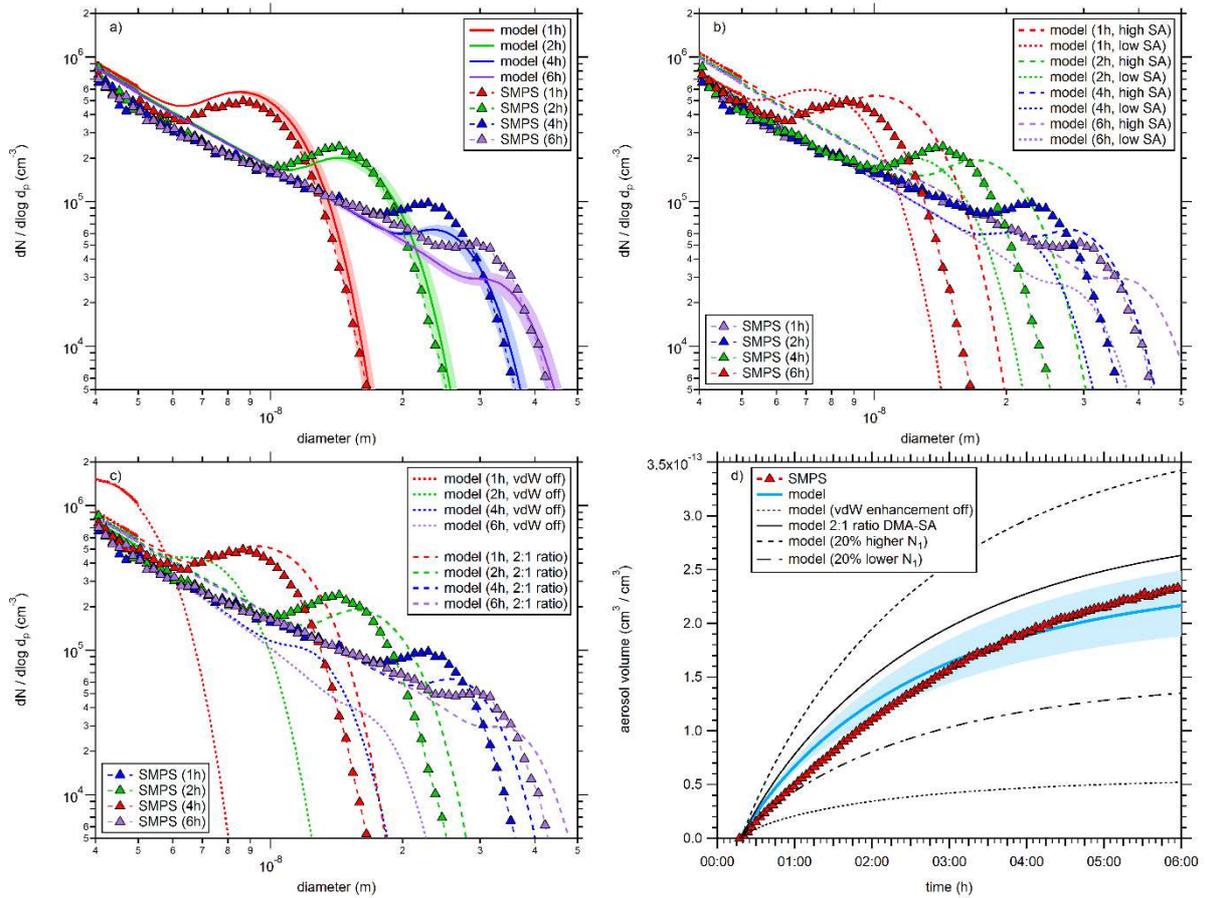
	kinetic model	model with evaporation rates
used for	Fig. 1, Fig. 2, Fig. 3 upper panel (black lines)	Fig. 3 upper panel (colored lines), Fig. 3 lower panel
described in	section 2.4	section 2.5, Appendix A
evaporation rates	all zero	$k_{e,A1B1} = 0.1 \text{ s}^{-1}$ $k_{e,A3B1} = 1 \text{ s}^{-1}$ $k_{e,A3B2} = 1 \text{ s}^{-1}$ $(k_{e,A4B1} = \infty \text{ s}^{-1})$ all others zero

1244



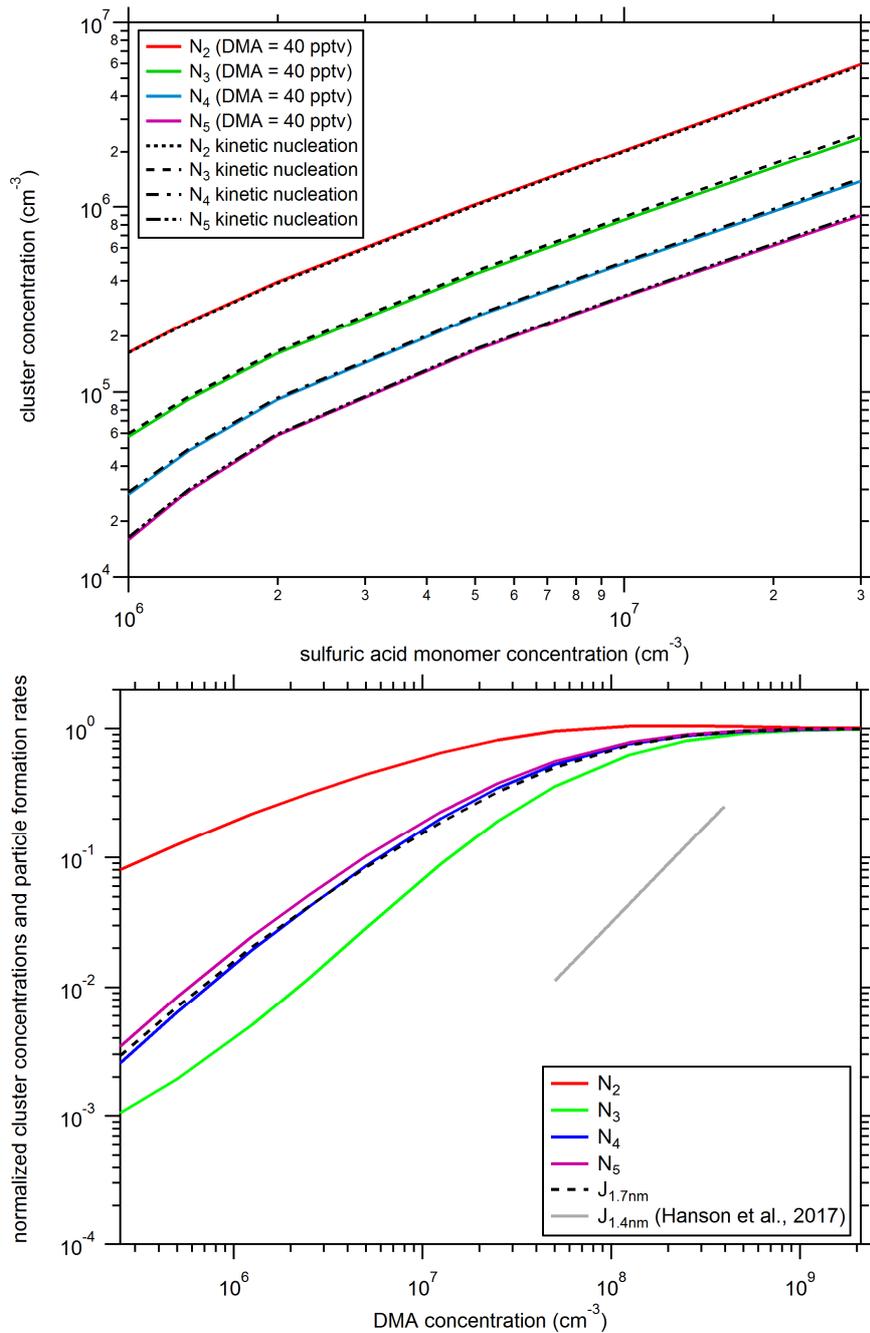
1245
 1246
 1247
 1248
 1249
 1250
 1251
 1252
 1253
 1254
 1255
 1256
 1257
 1258

Fig. 1. Comparison between experimental and theoretical particle formation rates at different sizes (mainly at mobility diameters of 1.7 nm and 4.3 nm). The lines indicate calculated particle formation rates from the collision-controlled aerosol model described in section 2.4 for CLOUD chamber conditions. The shaded regions show the model uncertainties when using an error of $\pm 20\%$ for the wall loss coefficient (C_w , see equation (2)). The open red symbols show previously published CLOUD7 data for the sulfuric acid-dimethylamine-water system (Almeida et al., 2013), while the blue symbols show the rates derived from SMPS size distribution measurements (this study). The data shown by the closed red symbols were derived with the method introduced by Kürten et al. (2015a) by extrapolating the SMPS data starting at 4.3 nm. The black lines show the calculated formation rates from the ACDC model for a mobility diameter of 1.2 to 1.4 nm (Almeida et al., 2013). Equation (10) from Hanson et al. (2017) is used to generate the green line.



1259
1260

1261 **Fig. 2.** Comparison between simulated and measured particle size distributions for one
 1262 experiment (CLOUD7, run 1036.01). The comparison is shown for four different times (1h, 2h,
 1263 4h and 6h) after the start of the experiment (panels a, b and c). Panel d shows a comparison
 1264 between modeled and measured aerosol volume as a function of time. The shaded regions in
 1265 panel a show the model uncertainties when using an error of $\pm 20\%$ for the wall loss coefficient
 1266 (C_w , see equation (2)). Panel b shows the change in the size distributions when the sulfuric acid
 1267 monomer concentration is varied by $\pm 20\%$. The effect of van der Waals forces on the size
 1268 distribution is shown in panel c along with the assumption that particles grow by the addition
 1269 of 2 DMA and 1 sulfuric acid molecule (2:1 ratio instead of 1:1 ratio). See text for further
 1270 details.



1271
 1272
 1273
 1274
 1275
 1276
 1277
 1278
 1279
 1280
 1281
 1282
 1283

Fig. 3. Upper panel: Comparison of modeled cluster ($N_2 =$ dimer, $N_3 =$ trimer, $N_4 =$ tetramer and $N_5 =$ pentamer) concentrations using different scenarios. The dashed black lines use the collision-controlled nucleation scheme with all evaporation rates set to zero (section 2.4); while the colored solid lines are calculated based on the model from section 2.5 with a dimethylamine (DMA) mixing ratio of 40 pptv ($1 \times 10^9 \text{ cm}^{-3}$), which was the average mixing ratio during the CLOUD7 campaign. Lower panel: Variation in modeled cluster concentration and $J_{1.7\text{nm}}$ as a function of the dimethylamine mixing ratio. The data were normalized to the values from the collision-controlled limit calculation (upper panel). For the calculations, a sulfuric acid monomer concentration of $N_1 = 5 \times 10^6 \text{ cm}^{-3}$ was used. An expression from Hanson et al. (2017) to calculate NPF rates as a function of DMA is shown by the grey line. See text for further details.