



- 1 Characterization of free amino acids, bacteria and fungi in size-
- 2 segregated atmospheric aerosols in boreal forest: seasonal patterns,

# 3 abundances and size distributions

- 4
- 5 Aku Helin<sup>1</sup>, Outi-Maaria Sietiö<sup>2</sup>, Jussi Heinonsalo<sup>2</sup>, Jaana Bäck<sup>3</sup>, Marja-Liisa Riekkola<sup>1</sup> and Jevgeni
- 6 Parshintsev<sup>1</sup>
- 7 <sup>1</sup>Department of Chemistry, P.O. Box 55, FI-00014 University of Helsinki, Finland
- <sup>2</sup> Division of Microbiology and Biotechnology, Department of Food and Environmental Sciences, University of Helsinki,
   P.O. Box 56, FI-00014, Finland
- 10 <sup>3</sup> Department of Forest Sciences, University of Helsinki, P.O. Box 27, FI-00014, Finland
- 11 Correspondence to: Jevgeni Parshintsev (evgeny.parshintsev@helsinki.fi)
- 12
- 13
- 14
- 15 Abstract
- 16
- 17 Bioaerosols are ubiquitous in the atmosphere and constitute  $\sim$ 30% of atmospheric aerosol particle mass in sizes >1 µm. 18 Bioaerosol components, such as bacteria, fungi and pollen, may affect the climate by acting as could-active particles, thus 19 having an effect on cloud and precipitation formation processes. In this study, size-segregated aerosol samples (<1.0, 1-20 2.5, 2.5-10 and >10 µm) were collected in boreal forest (Hyytiälä, Finland) during one year and analyzed for free amino 21 acids (FAAs), DNA concentration and microorganism (bacteria, Pseudomonas and fungi). Measurements were performed 22 using tandem mass spectrometry, spectrophotometry and qPCR, respectively. Meteorological parameters and statistical 23 analysis were used to study their atmospheric implication for results. Distinct annual patterns of bioaerosol components 24 were observed, late spring and autumn being seasons of dominant occurrence. Elevated abundances of FAAs and bacteria 25 were observed during the local pollen season, whereas fungi were observed at highest level during autumn. 26 Meteorological parameters, such as air and soil temperature, radiation and rainfall were observed to possess close 27 relationship with bioaerosol abundances on an annual scale. 28 29 30 31 32 33 34
- 35
- 36
- 37
- 38
- 39





## 40 1 Introduction

41

42 Bioaerosols are emitted directly from the biosphere into the atmosphere (Després et al., 2012;Fröhlich-Nowoisky et al., 43 2016). Bioaerosol are released from multiple sources, such as soil, vegetation and oceans, and they include e.g. pollen, 44 plant fragments, spores, bacteria, algae and viruses. In recent years, the abundance and dispersal of microorganisms in 45 the atmosphere has attracted more and more interest, mainly due to the underestimation of their abundance and their 46 possible atmospheric impact (Jaenicke, 2005;Morris et al., 2011;Deguillaume et al., 2008;Burrows et al., 2009b;Estillore 47 et al., 2016). It is estimated that globally bioaerosols constitute  $\sim 30\%$  of aerosol particles mass in the particles sized >1 48 μm in urban and rural air (Fröhlich-Nowoisky et al., 2016). This fraction can be much higher (~80%) in the tropical forest 49 areas (Pöschl et al., 2010;Elbert et al., 2007), and interestingly even up to 65% at boreal forest during pollen season (Manninen et al., 2014). These high percentages provide the basis for assumptions that bioaerosols may play an important 50 51 role in the atmosphere by affecting cloud and precipitation formation processes by acting as cloud and ice nuclei (Huffman 52 et al., 2013;Burrows et al., 2009b;Burrows et al., 2009a; Després et al., 2012; Fröhlich-Nowoisky et al., 2016). Thus, to 53 clarify their atmospheric transport and ecosystem interactions, bioaerosols' chemical and microbial constituents need 54 more characterization and identification.

55

56 Bioaerosols have been studied by using a variety of techniques (Georgakopoulos et al., 2008), but cultivation and 57 microscopy have been frequently employed for the elucidation of microorganisms in aerosols (Manninen et al., 58 2014;Després et al., 2012). Nowadays, molecular genetic analysis techniques, such as quantitative polymerase chain 59 reaction (qPCR) or next-generation sequencing (NGS), have gained ground, because they provide information not only 60 on the viable and cultivable cells, but also on uncultivable, dead or fragments of plant and animal cells (Després et al., 61 2012). Alternative approach to determine particles of biological origin is based on chemical tracer techniques (Bauer et 62 al., 2008;Gosselin et al., 2016;Hock et al., 2008;Schneider et al., 2011;Zhang et al., 2010;Staton et al., 2015;Rathnayake 63 et al., 2017). The benefit of chemical tracers, such as carbohydrates, proteins/amino acids and lipids, is their applicability 64 to quantitative analysis, while their main disadvantage is disability to identify different biological species. Specific tracers 65 have been used to estimate the total amount of bioaerosols in the atmosphere (Gosselin et al., 2016;Hock et al., 66 2008;Schneider et al., 2011;Zhu et al., 2015).

67

Biologically-derived amino acids are non-volatile and hygroscopic compounds that are mainly found in the condensed phases in the atmosphere (Matos et al., 2016;Samy et al., 2013;Zhang and Anastasio, 2003). They are present in aerosols either in the combined form (proteins and peptides) or as free amino acids (FAAs), and they are emitted either from biogenic sources or formed from the degradation of proteinaceous material (Milne and Zika, 1993;Matos et al., 2016). Due to their close relationship with bioaerosols, amino acids can be used as biomarker for an overall estimation of biomass in aerosols (Hock et al., 2008;Schneider et al., 2011).

74

75 In this study size-segregated aerosol samples were collected in boreal forest during one year period. Aerosol samples 76 were analyzed for DNA concentration, microorganism-DNA (bacteria, *Pseudomonas* and fungi) and FAAs. The annual 77 concentration and size distribution variation of microorganisms and FAAs were investigated in detail in order to

vunderstand their potential sources. Correlation study between the bioaerosol components and meteorological parameters





| 79  | was carried out. Our objective was to gain further information about the abundances of bioaerosols in different particle       |
|-----|--|
| 80  | size fractions at the boreal forest region and to understand better their biosphere-atmosphere interactions.                   |
| 81  |  |
| 82  | 2 Experimental section   |
| 83  |  |
| 84  | 2.1 Materials and reagents   |
| 85  |  |
| 86  | Detailed information on materials and reagents is presented in the Supporting Information (SI).                                |
| 87  |  |
| 88  | 2.2 Aerosol sampling   |
| 89  |  |
| 90  | The aerosol samples were collected in a Scots pine forest between February and October 2014 at the SMEAR II station            |
| 91  | (Station for Measuring Forest Ecosystem-Atmosphere Relations) in Hyytiälä, Finland (Hari and Kulmala, 2005). A Dekati          |
| 92  | PM10-impactor was used for the sampling of four particle size fractions (<1.0, 1-2.5, 2.5-10 and >10 $\mu$ m) below the        |
| 93  | canopy. The sampling flow rate was on average 30 L/min and the collection time was from one to four days (sampling             |
| 94  | volume 76-144 m <sup>3</sup> ). The collection filters were 25 mm polycarbonate membranes (Whatman Nuclepore) for the three    |
| 95  | largest particle size fractions. To prevent particles bouncing, membranes were smeared with diluted Apiezon L vacuum           |
| 96  | grease. The smallest size fraction (<1.0 $\mu$ m) was collected on a 47 mm Teflon filter (Gelman Sciences) with 2 $\mu$ m pore |
| 97  | size. After sampling, the filters were dried and weighted. The sampling procedure is described in more details in Laakso       |
| 98  | et al. (2003). After gravimetric analysis, the filters were placed inside a closed polystyrene petri dish, covered with        |
| 99  | aluminum foil and stored at -20 °C.  |
| 100 |  |
| 101 | Two sets of aerosol samples were collected in consecutive days, so that the sampling dates were close to each other (Table     |
| 102 | S1). The first set of aerosol samples were analyzed for DNA and microorganisms (set A) and the other set for free amino        |
| 103 | acids (set B). For molecular biological analysis only particles larger than 1 µm were selected because smaller fractions       |
| 104 | were suspected to contain only virus particles and fragmented DNA.   |
| 105 |  |
| 106 | 2.3 Determination of amino acids   |
| 107 |  |
| 108 | After ultrasonic assisted extraction, free amino acids were determined by liquid chromatography-tandem mass                    |
| 109 | spectrometry. Detailed information about the validation and quality control of the analytical method is presented in the       |
| 110 | Supporting Information, Figure S1 and Tables S2-S4.  |
| 111 |  |
| 112 | 2.4 Extraction of DINA   |
| 113 | Total nucleic acide ware extracted from the collection filters with a commercial DNA extraction kit (BeweeWeter DNA            |
| 114 | Isolation Kit MoBio I shoratories IISA) with slight modification (see Supporting Information). The DNA concentration           |
| 116 | and purity was measured spectrophotometrically at 260 nm and 280 nm with NanoDrop ND-1000 (Thermo Scientific                   |
| 110 | and purity was inclusived spectrophotonetrearly at 200 million with real purity in 200 (1101110) Selentine,                    |

- 117 USA). The DNA samples were stored at -20  $^{\circ}\mathrm{C}$  prior to qPCR.
- 118





- 119 2.5 DNA amplification
- 120
- 121 The bacterial and fungal DNA amounts of the filter samples were quantified with qPCR using target specific primers
- 122 pairs, Eub338F and Eub518R, and FF390 and FR1 (Table S5) (Fierer et al., 2005; Vainio and Hantula, 2000). In addition,
- 123 genus specific primers, Eub338F and PseudoR, were utilized to detect the bacteria belonging to the genus *Pseudomonas*
- 124 (Purohit et al., 2003).
- 125

126 In the bacterial and Pseudomonas specific qPCR reactions, standard curves were generated with DNA extracted 127 Pseudomonas fluorescens H-27 (Hambi culture collection, University of Helsinki) and for the fungal specific qPCR, the 128 DNA from the newly whole genome sequenced Phlebia radiata FBCC43 (genome size 40.92 Mb, FBCC culture 129 collection, University of Helsinki) was used (Kuuskeri et al., 2016). When converting the copy number of the samples 130 into bacterial cells (/colony forming units), all bacteria were assumed to contain average of three gene copies of ribosomal 131 16S DNA per cell, and bacteria belonging to genus Pseudomonas was assumed to contain five gene copies of 16S rDNA 132 in a cell (Stoddard et al., 2014). According to Fröhlich-Nowoisky et al. (2009), majority of the fungal DNA detected 133 during summer and autumn are from fungal species belonging to Basidiomycota. Based on previous results (Fröhlich-134 Nowoisky et al., 2012; Fröhlich-Nowoisky et al., 2009), we have assumed that most the fungal DNA collected on our 135 filters is also derived from basidiomycete fungi. Because only few 18S rDNA-regions of fungi are fully annotated, we 136 have selected the *P. radiata* as a model basidiomycete fungus and calculated the gene copy numbers to fungal cells based 137 on the assumption that all fungi in this experiment have approximately the same amount of 18S rDNA gene copies as this 138 fungus has (Kuuskeri et al., 2016).

- 139
- 140 2.6 Additional background data and back-trajectory analysis
- 141

142 Meteorological variables, gas fluxes and atmospheric gases are continuously measured at the SMEAR II and the data is 143 available from AVAA-portal (Junninen et al., 2009). Half-hourly averaged data from the portal was further averaged 144 according to each sampling time period by using arithmetic mean. These averaged values were used in statistical analyses. 145 The selected variables were air (AT) and soil surface temperature (SST), soil surface water content (SSWC), wind speed 146 (WS) and direction (WD), gross primary production (GPP), CO, CO2 and ozone concentration, photosynthetically active 147 radiation (PAR), UV-A and UV-B radiation and relative humidity (RH). In addition, rainfall was averaged according to 148 different time periods as follows: 72 h before sampling (BSR), during sampling (DSR) and 168 h after sampling (ASR). 149 150 Backward air mass trajectories were calculated using the HYSPLIT (Hybrid Single-Paricle Lagrangian Integrated 151 Trajectory) transport and dispersion model from NOAA Air Resources Laboratory to estimate the origin and transport

152 route of aerosol particles (Draxler and Hess, 1998;Stein et al., 2015;Rolph, 2003). For the calculations, meteorological

- data from the Global Data Assimilation System (GDAS, 1 degree, global, 2006-present) was used. The backward air mass
- 154 trajectories were modelled 48 hours back time during the whole period of sampling, using a resolution of 12-hours.
- 155

156 2.7 Statistical analysis

157





158 For analyzing the importance of the seasonality and the aerosol filter size, permutational multivariate analysis of variance 159 (PERMANOVA) was performed for the presence/absence transformed amino acid data. The PERMANOVA was 160 performed with the adonis-function of the vegan package (Oksanen et al., 2016) with 999 permutations, and the amino 161 acid data was set as response variables while sampling month and aerosol filter size were set as explanatory variables. 162 One-way analysis of variance (ANOVA) was performed to the fungal, bacterial and Pseudomonas gene copy numbers 163 individually with the aov-funcion from the stats package (R Core Team, 2016) in order to study the effect of seasonality 164 and aerosol filter size. Prior to performing ANOVA the normal distribution of each residual was checked individually 165 with shapiro.test-function, and all the gene copy number data was log-normalized. The linear correlations between the 166 FAA and microorganisms concentrations with meteorological variables were calculated using the rcorr-function of the 167 Hmisc package (Harrell, 2016) and visualized with the corrplot package (Wei and Simko, 2016) showing only 168 correlations with statistical significance ( $p \le 0.05$ ).

169

170

## 171 3 Results and Discussion

172

173 3.1 General characteristics

174

175 The average concentration and range of each component measured in different particle size fractions at the boreal forest 176 site is presented in Table 1 and Table S6. The lowest concentration levels of FAA and total DNA were measured during 177 winter and highest during late spring (Figures 1 and S2). The total amount of extracted DNA was highest from May to 178 June (up to 48 ng/m<sup>3</sup>) and from September to October (up to 14 ng/m<sup>3</sup>). The FAA concentration in total particles was 179 highest in late May (up to 751 ng/m<sup>3</sup>) and in September (up to 35 ng/m<sup>3</sup>). A common trend in both total DNA and FAA 180 concentrations was a maximum in spring and a secondary smaller maximum in autumn. Similarly, the highest PM 181 concentration in total particles was recorded during late spring and autumn (Figure S3); a pattern which has also been 182 previously observed at the same location (Laakso et al., 2003; Manninen et al., 2014). The measured DNA concentrations 183 agreed with those obtained for PM2.5 samples at a rural mountain site in Germany, where the DNA concentration was in 184 the range 1.7-4.2 ng/m<sup>3</sup> (Després et al., 2007). Also the observed FAA concentrations were in the same order of magnitude 185 as those observed in other rural and semi-urban locations (Zhang and Anastasio, 2003;Zhang et al., 2002;Samy et al., 186 2011;Samy et al., 2013). For example, at two rural sites in US the concentration of FAAs in PM<sub>2.5</sub> samples was measured 187 to be 22±9 ng/m<sup>3</sup> during summer (Samy et al., 2011), and 59±49 ng/m<sup>3</sup> (range 9-236 ng/m<sup>3</sup>) during one year study period 188 (Zhang and Anastasio, 2003). To our knowledge, there are no studies covering DNA and FAA abundances at the boreal 189 forest region, thus comparison to previous results is not feasible.

190

Similar to the annual trend of total DNA and FAAs, the lowest concentration levels of microorganisms were detected during winter and highest during spring and autumn. However, bacterial and fungal DNA reached their maximum levels at different seasons; bacteria peaking in late spring (Figure 2a) and fungal DNA in late summer and autumn (Figure 3). The bacterial DNA abundance in total particles was lower than 900 cells/m<sup>3</sup> during winter and early spring, whereas in late spring the concentration peaked at 58731 cells/m<sup>3</sup>. During summer and autumn, the concentration of bacterial DNA was one magnitude lower. In contrast to the total number of bacteria, the highest amounts (>200 cells/m<sup>3</sup>) of *Pseudomonas* DNA were measured in late May, late June and October (Figure 2b). In rest of times, the concentration of bacteria





198 belonging to the genus Pseudomonas was much lower, and the highest concentration levels in total particles rarely 199 exceeded 50 cells/m<sup>3</sup>. The fungal abundance was lower than 300 cells/m<sup>3</sup> during winter and early spring, whereas in late 200 spring and summer the concentrations started to increase and generally exceeded 6000 cells/m<sup>3</sup> in total particles. The 201 highest amounts of fungal DNA were measured in late June (30149 cells/m<sup>3</sup>), in August (55839 cells/m<sup>3</sup>) and in early 202 October (35050 cells/m<sup>3</sup>). In general, the concentration levels of microorganisms vary seasonally and geographically, but 203 are estimated to be in the level of  $\sim 10^4$ - $10^5$  m<sup>-3</sup> and  $\sim 10^4$ - $10^5$  m<sup>-3</sup> for bacterial cells and fungal spores, respectively 204 (Burrows et al., 2009b;Després et al., 2012;Spracklen and Heald, 2014). Our observations are consistent with the common 205 trend, when considering that low concentration levels are typically observed at rural locations.

- 206
- 207
- 208



209

Figure 1. Free amino acid concentrations over the sampling period in different particle size fractions (date format isdd.mm.yyyy.). Note the different y-axis scales in panels.







212

213 Figure 2. a) The amount of all bacterial cells over the sampling period, and b) cells of bacteria belonging to genus

- 214 *Pseudomonas*, detected with qPCR from the aerosol filters, specified by the particle size fractions (date format is
- 215 dd.mm.yyyy.).
- 216 217



218

219 Figure 3. The amount of fungal cells and spores over the sampling period, detected with qPCR from the filters of

- 220 different particle size fractions (date format is dd.mm.yyyy.).
- 221

222 3.2 Seasonal variation and size distribution of FAAs

223

In winter the amino acids were mainly accumulated in the particle size fraction <1  $\mu$ m (~74%, Figure S4a), whereas during spring the amino acids were mainly present in the largest particle size fraction >10  $\mu$ m (~77%). The highest amount

226 of FAAs were observed during late spring and early summer (Figure 1), when the local pollen season plumed. During

227 summer, the concentrations of FAAs were relatively constant and seemingly FAAs evenly distributed in all the size





fractions (Figure S4a). In autumn they were mainly present in the particles below 10 μm, and accumulated especially in the size fraction 2.5-10 μm (~51%). During autumn, there was a slight increase in the total FAA concentrations mainly due to the increase in the size fraction 2.5-10 μm (Figure 1).

231

232 Similarly, in the size distribution of individual amino acids some clear tendencies were observed (Figure S5). The results 233 from the PERMANOVA analysis indicated that seasonal variation explained the observed variation across the whole 234 FAA data with statistical significance (p<0.001). In addition, the size of the aerosol filter was statistically significant 235 factor for explaining the observed variation across the whole amino acid data (PERMANOVA, p<0.001). Glycine was 236 the most dominant amino acid in the size fraction  $<1 \,\mu$ m, whereas in the larger size fractions its relative abundance was 237 much smaller (Figure S5). On average, glycine accounted 59% of the total FAAs in particles <1 µm, followed by alanine 238 (10%), aspartic acid (9%) and glutamine (9%). In the size fraction 1-2.5 µm, glutamine (42%), glutamic acid (16%) and 239 arginine (15%) were on average the most abundant amino acids. Similarly, in the size fraction 2.5-10 µm glutamine 240 (29%), arginine (15%) and glutamic acid (12%) were the dominant ones, accompanied by proline (7%), serine (7%) and 241 alanine (7%). In the largest size fraction >10 µm proline (36%), arginine (21%), glutamine (10%) and glutamic acid (9%) 242 were the most dominant ones. Our results are in agreement with several studies demonstrating that glycine, arginine, 243 serine, alanine, proline and glutamine are the most commonly abundant FAAs in aerosols (Barbaro et al., 2011;Barbaro 244 et al., 2015;Samy et al., 2011;Samy et al., 2013;Scalabrin et al., 2012;Violaki and Mihalopoulos, 2010;Zhang and 245 Anastasio, 2003), although the relative abundances vary depending on location, season and particle size fraction (Matos 246 et al., 2016).

247

248 3.3 Seasonal variation and size distribution of microorganisms

249

250 A distinct monthly variation in bacterial DNA abundance in different filter size fractions was observed (ANOVA, p<0.01). 251 During early spring, bacterial DNA was accumulated mainly in the size fraction >10 µm (~64%, Figure S4b). Oppositely, 252 in late spring the highest numbers of bacteria were detected in the size fraction 1-2.5 µm, whereas during summer, highest 253 numbers of bacteria were discovered in the size fraction 2.5-10 µm (Figure 2a). During summer, bacteria was mainly 254 present in the size faction 2.5-10 µm, which covered on average ~57% of the total bacteria amount. In general, most of 255 the bacteria were observed in the size fractions 1-2.5  $\mu$ m and 2.5-10  $\mu$ m, and less bacterial cells were found in the size 256 fraction >10 µm. However, on average the abundances did not differ significantly in different size fractions (ANOVA, 257 p=0.494). The size distribution observed is in line with that of Burrows et al. (2009b) who summarized that the median 258 aerodynamic diameter of particles containing bacteria is 4 µm at continental sites.

259

260 The abundance of *Pseudomonas* varied with statistically significance between different months (ANOVA, p<0.05) and 261 between different size fractions (ANOVA, p<0.001). Pseudomonas was mainly present (~70%) in the smallest particle 262 size fraction 1-2.5 µm throughout the year (Figures 2b and S4c). The accumulation of Pseudomonas in the smallest size 263 fraction is in line with the small aerodynamic size (<1 µm) of common Pseudomonas species (Chang et al., 2013;Möhler 264 et al., 2008; Pietsch et al., 2015). Overall, the relative amount of Pseudomonas from all bacteria was highest in the size 265 fraction 1-2.5 µm (Figure S6), being highest in late winter (28%), in early spring (15%) and in autumn (23%). 266 Interestingly, in February the relative amount of *Pseudomonas* DNA of the total bacterial DNA was 17% in total particles 267 (Figure S6).





# 268

269 The amount of fungal DNA detected in different size fractions varied with statistical significance (ANOVA, p<0.05). The 270 fungal DNA was predominantly observed in the particle size fraction 2.5-10 µm, which accounted on average ~58% of 271 the total fungal concentration. The monthly variation in fungal DNA abundance was also statistically significant 272 (ANOVA, p<0.001). During May and early June, the fungal DNA was mainly accumulated in the size fraction 1-2.5 µm 273 (~67%, Figure S4d). In late summer, fungal DNA was detected at its highest level in the size fraction 2.5-10 µm (Figure 274 3), and in overall ~66% was accumulated in this size fraction during the summer months. Fungal spores are frequently 275 observed in aerosols in the size range of 2-10 µm (Després et al., 2012). Our observations are consistent with these 276 literature values. 277 278 3.4 Overview of meteorological factors and sources 279 280 The effect of local meteorological factors on the concentration levels of FAAs and microorganism was studied by means 281 of linear regression analysis and a summary of results is shown in Figure 4. Our results reflect rather long-term seasonal 282 effects (more details in Supporting Information), due to the time resolution of sampling. Thus, as expected, the key 283 meteorological factors explaining the observed concentration levels were air temperature, soil temperature, gross primary 284 production and radiation (Figures S7 and S8). In general, we believe that the positive correlation observed between FAAs 285 and microorganism abundances with the before mentioned meteorological variables are closely related to growing season 286 and seasonality, rather than being a decisive effect. These observations are in good agreement with previously reported 287 tendencies covering microorganism and different meteorological factors (Jones and Harrison, 2004; Burrows et al., 288 2009b;Lighthart, 2000). Soil and vegetation have been previously suggested to be the predominant sources of 289 microorganisms in the atmosphere (Bowers et al., 2013), and our correlation results confirm these sources as discussed 290 below.



291

Figure 4. Summary of Spearman correlation results (total particles, p<0.05). The colour scale indicates positive/negative</li>
correlation. Abbreviations: AT-air temperature; SST-soil surface temperature; SSWC-soil surface water content; WSwind speed; WD-wind direction; GPP-gross primary production; CO; CO2; ozone concentration; PAR-photosynthetically
active radiation; UV-A and UV-B radiation; RH-relative humidity; BSR-rainfall 72 h before sampling; DSR-rainfall
during sampling; ASR-rainfall 168 h after sampling.

297

As could be expected (Manninen et al., 2014;Schumacher et al., 2013), the lowest concentration levels of FAAs and microorganisms were detected during winter, when the air and soil temperatures were below 0 °C and the ground was covered by snow. In February, particles below 2.5 μm accounted for 89% of the total FAA concentration. In these samples, glycine and alanine were the dominant amino acids. Glycine and alanine have been exploited as markers for long-range





302 transport aerosols due to their relatively low reactivity and long half-lives in the atmospheric condensed phases (Barbaro 303 et al., 2011; Barbaro et al., 2015; Scalabrin et al., 2012; Samy et al., 2013). As the concentration levels of microorganisms 304 were low during winter (Figures 2 and 3), it is likely that the observed FAAs originated from distant sources. However, 305 FAAs and particularly glycine have been also associated with biomass burning emissions (Samy et al., 2013; Violaki and 306 Mihalopoulos, 2010;Zangrando et al., 2016). In Finland wood-burning, that is typical domestic heating system in winter, 307 is the most presumable source of over one third of the PM2.5 emissions (Laakso et al., 2003;Saarnio et al., 2012). The 308 relative abundance of glycine in the size fraction 1-2.5 µm was 85% during winter, whereas in other seasons its 309 contribution was negligible. Thus, the observed glycine concentrations might be associated with wood-burning emissions, 310 coming from either local or distant sources. This conclusion is supported by the observed positive correlation between 311 glycine with CO concentration in the size fractions  $<1 \mu m$  (R=0.54, p<0.05, Figure S8) and 1-2.5  $\mu m$  (R=0.80, p<0.001, 312 Figure S8). Furthermore, only glycine among the FAAs showed positive correlation with PM concentration in the size 313 fraction 1-2.5 µm (R=0.69, p<0.01, data not shown).

314

315 FAA and bacteria concentrations increased during the spring recovery and seemed to be influenced by the increase in 316 local biological activity. During early spring, FAAs and bacteria were mainly present in the largest size fractions (Figures 317 S4a and S4b), which may indicate the presence of pollen in aerosols, since microorganisms might be present on the 318 surface of pollen grains (Puc, 2003). Further, Manninen et al. (2014) observed pollen to be present at SMEAR II station 319 already before the local pollen season, most possible due to long-range transport from central Europe, where growing 320 season had already started. Interestingly though, during our sampling periods the back-trajectory analysis results showed 321 air-masses arriving primarily from Scandinavia and North Europe instead of central Europe. This could imply that the 322 larger sized particles are originating locally from some early pollinator rather than being long-range transported. In 323 addition, we observed an interesting phenomenon in early March when all FAAs, bacterial DNA and fungal DNA were 324 present mainly in the size fraction  $>10 \,\mu$ m. Then the corresponding relative abundances of their total amounts were 73%, 325 68% and 92%, respectively. During March sampling, the air temperature was almost constantly above zero degrees for 326 the first time in our campaign. The small increase observed in FAA and microorganism concentrations during this time 327 might indicate that a burst in bioaerosols emissions can occur when the temperature increases and snow starts to melt 328 revealing the ground vegetation and decomposed leaf litter for the first time after winter. Similar observations have been 329 previously made at the same site by Schumacher et al. (2013) in a study covering fluorescent biological aerosol particles 330 (FBAP), although they detected an increase in FBAP concentrations in November after first snowfall and snowmelt event. 331

332 The concentrations of FAAs and bacteria reached their maximum levels during the local pollen season peak in May 333 (Figures 1 and 2). Elevated amino acid concentrations have been observed during spring time also in other locations and 334 suspected to be influenced by pollen (Barbaro et al., 2011; Zhang et al., 2002). In our study, especially proline and arginine 335 concentrations increased during the pollen season peak. These amino acids have been shown to contribute significantly 336 to the total amino acid content of birch pollen (Ozler et al., 2009). In this study, the FAA concentrations increased in all 337 the size fractions (Figure 1), possibly due to rupture of relatively large ( $\sim$ 30 µm) pollen grains (Taylor et al., 2004;Visez 338 et al., 2015), supported by visibly yellow impactor plates in all the size fractions. During the peak in FAA and bacteria 339 concentrations, the air-masses were arriving in Hyytiälä from Eastern Europe and Baltic Sea. Due to the absence of 340 pollinating species in the sea region, the observed high concentration levels of FAAs and bacteria could be mainly 341 explained by local sources in the boreal forest. This is corroborated by the presence of methionine, cysteine and tryptophan





342 only in these aerosols samples. Particularly these amino acids are known to be highly reactive with short half-lives in the 343 atmospheric condensed phases (Scalabrin et al., 2012; Milne and Zika, 1993; McGregor and Anastasio, 2001). Thus, the 344 local pollen season likely explains our observations, although other factors may partly contribute. For example, high 345 ozone concentration and strong UV radiation were recorded during this time period, and in some studies ozone has been 346 demonstrated to promote the decomposition of protein and peptides into free amino acids increasing the FAA 347 concentrations (Samy et al., 2013; Mumford et al., 1972). Positive correlation was observed between FAA concentration 348 in the size fraction  $>10 \,\mu\text{m}$  with ozone (R=0.58, p<0.05, Figure S8), which was mainly attributed to coinciding peaks in 349 concentration levels during pollen season.

350

351 Opposite to FAA concentrations being highest in the largest size fractions, bacterial DNA abundance was highest in the 352 size fraction 1-2.5 µm during the pollen season peak. As spring proceeds, new foliage growth enables larger surface area 353 for epiphytic bacteria to occupy and grow on. Bacterial cells may be lifted from leaves into the air on pollen (Jones and 354 Harrison, 2004), which may explain the peak in bacterial abundances during this season. Bacteria may be present in the 355 air as individual cells, clump of cells or attached to other particles, such as pollen grains and leaf fragments. We propose 356 that during the local pollen season and under favourable meteorological conditions, it is possible to observe a significant 357 increase in bacterial concentrations. However, most likely the magnitude of this increase varies from year-to-year and is 358 also closely related to varying pollen abundances.

359

360 Although we assume that the elevated bacterial DNA abundances are mainly related to pollen and vegetation during 361 spring, other sources might contribute as well. For example, soil water content and bacteria concentration correlated 362 positively in the size fractions 1-2.5 µm (R=0.53, p<0.05, Figure S7) and >10 µm (R=0.57, p<0.05, Figure S7). After 363 snow melts, the soil moisture increase enhances the bacterial/microbial growth (Burrows et al., 2009b). When spring 364 proceeds and air and soil temperatures increase, the relatively dry soil surface layer might enable the dispersal of bacteria 365 via soil resuspension. Similarly, some studies indicate that soil-derived bacteria dominate during spring time (Rathnayake et al., 2017), while others indicate that soil sources dominate during late summer and fall (Bowers et al., 2013). In addition, 366 367 rainfall may promote the bacterial growth on vegetation surfaces, leading to increased population sizes, which may 368 become airborne following rainfall (Bigg et al., 2015). We observed positive correlation between bacteria concentration 369 in the size fraction 1-2.5 µm with rainfall recorded 72 h before the sampling (R=0.64, p<0.01, Figure S7). Based on 370 previous reports related to the relationship between rainfall and bioaerosols (Huffman et al., 2013; Prenni et al., 371 2013;Gosselin et al., 2016;Rathnayake et al., 2017;Morris et al., 2016;Bigg et al., 2015), our results corroborate the 372 positive effect of bacteria abundances following rainfall.

373

374 The abundance and relative size distribution of fungal DNA varied seasonally and started to increase in summer (Figure 375 3). The relative distribution of fungal DNA to different size fractions correlated with the relative humidity (R=-0.53, 376 P=0.035 for 1-2.5 µm and R=0.45, P=0.081 for 2.5-10 µm, data not shown). The lower the relative humidity was, the 377 more fungi were accumulated in the smallest size fraction. Since fungal cells are typically larger than 2.5 µm, the observed 378 fungal DNA in the 1-2.5 µm fractions during early summer is probably originated from the spores of moulds or ruptured 379 cells (Reponen et al., 2001). In addition, the size of fungal spores has been found to depend on the relative humidity, i.e. 380 higher the relative humidity is the larger the spores are (Reponen et al., 1996; Madsen, 2012). The spores of basidiomycete 381 fungi are in general larger than those of ascomycete (Reponen et al., 2001; Manninen et al., 2014; Hussein et al.,





2013;Fröhlich-Nowoisky et al., 2012), and the basidiomycete fungi are known to sporulate mainly during autumn when the relative humidity is high (Kauserud et al., 2011). Further, in the previous study by Manninen et al. (2014) the spores of Basidiomycota dominated the phylum-level distribution of fungal spores in the autumn. Our results where fungal DNA accumulated in the 2.5-10 µm size fraction during autumn are in agreement with this study. Consequently, the annual size

distribution of fungi can be expected to be similar from year-to-year in the boreal forest.

386 387

388 Opposite to the trend of fungi, FAA and bacterial DNA concentrations decreased after the spring pollen season peak. It 389 is noteworthy that both bacterial and fungal DNA were accumulated during the mid-summer to early fall in the particle 390 size fraction 2.5-10 µm (Figure S4b and S4d). In summer, the air masses arrived Hyytiälä mainly from Scandinavia and 391 Baltic Sea regions. In early August the air masses reached Hyytiälä from a large forest fire region in Sweden. Interestingly 392 derived from our samples, glycine was detected in relatively high concentrations ( $\sim 2.9 \text{ ng/m}^3$ ) in the size fraction <1  $\mu$ m. 393 This observation supports the hypothesis that the presence of glycine is partially related to biomass burning emissions 394 and long-range transport. However, no distinct effects on the origin of air-masses were seen in the levels of FAAs or 395 microorganisms. Although long-range transport cannot be completely ruled out, the accumulation of bacteria and fungi 396 in the particle size fraction 2.5-10 µm suggest that primarily local forest or nearby sources affect the concentration levels 397 of microorganisms during summer. For example, leaf-associated bacteria are known to contribute to the total bacterial 398 amount during summer (Bowers et al., 2013). Statistically significant correlation was observed between gross primary 399 production (R=0.79, P<0.001, Figure S7) and photosynthetically active radiation (R=0.69, P<0.01, Figure S7) with 400 bacteria in the size fraction 2.5-10 µm, possibly indicating the importance of vegetation especially during summer and 401 growing season.

402

403 The abundance of *Pseudomonas* was detected at its highest level in early October in the size fraction 1-2.5 µm, when it 404 was 2-fold higher than in other months (Figure 2b). It is suggested that an increase in the Pseudomonas syringae 405 population size on vegetation may occur during or after rainfall (Hirano et al., 1996; Bigg et al., 2015; Morris et al., 2016). 406 Even though no correlation was observed between average rainfall recorded before or during sampling with Pseudomonas 407 on an annual scale, we identified interesting separate individual rainfall events. The maximum single rainfall event was 408 recorded within 72 h before the above mentioned early October sampling (Figure S9). As far as we could tell, this was 409 the main exception in meteorological conditions during the period studied, i.e. the only factor potentially explaining the 410 observed increase in Pseudomonas concentration. Further, the back-trajectory analysis results showed air-masses arriving 411 from Norwegian Sea and North Sea via the Scandinavian Mountains, and to our knowledge these areas do not contain 412 sources that could explain the increase in Pseudomonas concentration. When considering the other maximum rainfall 413 events recorded before sampling, two out of three of the summer-autumn maximum events coincided with high 414 Pseudomonas concentrations (Figure S9). These findings corroborate the evidence of a potential relationship between 415 rainfall and Pseudomonas. However, the lack of statistically significant correlation suggests that other factors are 416 contributing to observed variation in Pseudomonas abundance.

417

418 Overall, considering the effect of rainfall on the levels of bacteria, fungi and FAAs, some interesting patterns were 419 observed. As mentioned earlier, the bacterial DNA concentration in size fraction 1-2.5  $\mu$ m correlated with rainfall 420 recorded prior to sampling. Further, we observed positive correlation between FAA concentration and rainfall during 421 sampling in the size fractions 1-2.5  $\mu$ m (R=0.53, p=0.0502, data not shown) and 2.5-10  $\mu$ m (R=0.70, p<0.01, Figure S8).





422 Interestingly, bacterial DNA and fungal DNA abundances correlated positively in the size fraction 2.5-10 µm with rainfall 423 recorded 168 h after sampling (R=0.63, p<0.01 and R=0.53, p<0.05, respectively; Figure S7). There is relatively strong 424 evidence that cloud-active particles larger than ~1 µm are biological in origin (Haga et al., 2014;Hassett et al., 2015;Mason 425 et al., 2016; Möhler et al., 2007), and in mixed-phase cloud conditions, bioaerosols may play an important role in triggering 426 rainfall. Collectively, the different correlations observed between rainfall and bioaerosol components suggest that a 427 potential feedback mechanism may persist at the boreal forest. This conclusion is based on the assumption that the 428 recorded rainfall events after sampling were produced at least partially on a local scale and that the meteorological factors 429 were favourable to formation of rainfall. We acknowledge, that the positive correlation observed between 430 microorganism's abundances with rainfall recorded following sampling, may be causal in nature. However, in light of the 431 recent findings, the possibility of a feedback mechanisms cannot be ignored (Bigg et al., 2015;Morris et al., 2016;Huffman 432 et al., 2013). Nonetheless, additional high time resolution and long-term measurements are needed to confirm the 433 observations presented in this research.

434

### 435 4 Conclusions

436

437 Considering the observations made in this and previous studies (Manninen et al., 2014;Schumacher et al., 2013), some 438 general conclusions related to bioaerosol abundances and size distribution at the boreal forest site can be drawn. 439 Correlation with meteorological parameters might indicate biosphere-atmosphere interactions through bioaerosols with 440 possible climate effects. In early spring, bioaerosol components are mainly accumulated in the size-fractions >10 µm. 441 According to our results, the spring pollen season has an impact on the pollen levels as well as on the bacterial abundances. 442 Elevated bioaerosol abundances occur during the pollen season (Manninen et al., 2014), and based on our estimation even 443 up to ~77% of total PM may be of biological origin (SI). Thus, the magnitude of biological cloud-active particles during 444 this period may be atmospherically relevant (Diehl et al., 2002;Diehl et al., 2001;Pummer et al., 2012;Pummer et al., 445 2015). The importance of rainfall was observed in this study as well as to some extent in previous studies conducted at 446 the same site (Manninen et al., 2014;Schumacher et al., 2013). Positive correlation was observed between bioaerosol 447 component abundances with rainfall recorded before and during sampling as well as with rainfall recorded after sampling. 448 During late summer and autumn, the accumulation of bioaerosols in the size fraction 2.5-10 µm was evident (Figure S4). 449 In autumn the relative amount of biomass in PM was estimated to be around ~10% (SI). Modelling studies have suggested 450 that microorganisms may play an important role in the hydrological cycle in the boreal region (Sesartic et al., 2012;Sesartic 451 et al., 2013). However, the magnitude of upward lifting of microorganisms remains to be solved, and thus our results are 452 preliminary in nature and need to be confirmed. 453

454

-77

# 455

# 456 Author contribution

457 A.Helin, O.-M. Sietiö, J. Heinonsalo, J. Bäck, M.-L. Riekkola, J. Parshintsev designed the experiments. A. Helin, O.-M.

458 Sietiö, J. Parshintsev carried them out. A. Helin and O.-M. Sietiö performed the statistical analysis. A. Helin and J.

- 459 Parshintsev prepared the manuscript with contributions from all co-authors.
- 460 **Competing interests**





461 The authors declare that they have no conflict of interest.

#### 462 Acknowledgements

- 463 The financial support of the Academy of Finland Center of Excellence program (project no 272041, JP, MLR, AH) and
- 464 research project no 292699 (OMS, JH) are gratefully acknowledged. Technical staff of the SMEAR II station are thanked
- 465 for their valuable help. Magnus Ehrnrooth foundation (JP) and University of Helsinki Doctoral Program in Microbiology
- 466 and Biotechnology (MBDP) (OMS) are thanked for support. Geoffroy Duporté is acknowledged for back-trajectory
- 467 analysis. Merck Life Science is thanked for providing ZIC-cHILIC columns.
- 468

### 469 References

- 470 Barbaro, E., Zangrando, R., Moret, I., Barbante, C., Cescon, P., and Gambaro, A.: Free amino acids in atmospheric 471 particulate matter of Venice, Italy, Atmos. Environ., 45, 5050-5057, 2011.
- 472 Barbaro, E., Zangrando, R., Vecchiato, M., Piazza, R., Cairns, W., Capodaglio, G., Barbante, C., and Gambaro, A.: Free 473 amino acids in Antarctic aerosol: potential markers for the evolution and fate of marine aerosol, Atmos. Chem. Phys., 15,
- 474 5457-5469, 2015.
- 475 Bauer, H., Claeys, M., Vermeylen, R., Schueller, E., Weinke, G., Berger, A., and Puxbaum, H.: Arabitol and mannitol as 476 tracers for the quantification of airborne fungal spores, Atmos. Environ., 42, 588-593, 2008.
- 477 Bigg, E. K., Soubeyrand, S., and Morris, C. E.: Persistent after-effects of heavy rain on concentrations of ice nuclei and
- 478 rainfall suggest a biological cause, Atmos. Chem. Phys., 15, 2313-2326, 2015.
- 479 Bowers, R. M., Clements, N., Emerson, J. B., Wiedinmyer, C., Hannigan, M. P., and Fierer, N.: Seasonal variability in 480 bacterial and fungal diversity of the near-surface atmosphere, Environ. Sci. Technol., 47, 12097-12106, 2013.
- 481 Burrows, S. M., Butler, T., Jöckel, P., Tost, H., Kerkweg, A., Pöschl, U., and Lawrence, M. G.: Bacteria in the global 482 atmosphere - Part 2: Modeling of emissions and transport between different ecosystems, Atmos. Chem. Phys., 9, 9281-483 9297, 2009a.
- 484 Burrows, S. M., Elbert, W., Lawrence, M. G., and Pöschl, U.: Bacteria in the global atmosphere - Part 1: Review and 485 synthesis of literature data for different ecosystems, Atmos. Chem. Phys., 9, 9263-9280, 2009b.
- 486 Chang, C. W., Li, S. Y., Huang, S. H., Huang, C. K., Chen, Y. Y., and Chen, C. C.: Effects of ultraviolet germicidal 487 irradiation and swirling motion on airborne Staphylococcus aureus, Pseudomonas aeruginosa and Legionella 488 pneumophila under various relative humidities, Indoor air, 23, 74-84, 2013.
- 489 Deguillaume, L., Leriche, M., Amato, P., Ariya, P., Delort, A., Pöschl, U., Chaumerliac, N., Bauer, H., Flossmann, A.,
- 490 and Morris, C.: Microbiology and atmospheric processes: chemical interactions of primary biological aerosols, 491 Biogeosciences Discussions, 5, 841-870, 2008.
- 492 Després, V., Nowoisky, J., Klose, M., Conrad, R., Andreae, M., and Pöschl, U.: Characterization of primary biogenic 493 aerosol particles in urban, rural, and high-alpine air by DNA sequence and restriction fragment analysis of ribosomal 494 RNA genes, Biogeosciences, 4, 1127-1141, 2007.
- 495 Després, V. R., Huffman, J. A., Burrows, S. M., Hoose, C., Safatov, A. S., Buryak, G., Fröhlich-Nowoisky, J., Elbert,
- 496 W., Andreae, M. O., and Pöschl, U.: Primary biological aerosol particles in the atmosphere: a review, Tellus B, 64, 2012. 497 Diehl, K., Quick, C., Matthias-Maser, S., Mitra, S. K., and Jaenicke, R.: The ice nucleating ability of pollen: Part I:
- 498 Laboratory studies in deposition and condensation freezing modes, Atmos. Res., 58, 75-87, 2001.
- 499 Diehl, K., Matthias-Maser, S., Jaenicke, R., and Mitra, S. K.: The ice nucleating ability of pollen:: Part II. Laboratory 500 studies in immersion and contact freezing modes, Atmos. Res., 61, 125-133, 2002.
- 501 Draxler, R. R., and Hess, G.: An overview of the HYSPLIT\_4 modelling system for trajectories, Aust. Meteorol. Mag., 502 47, 295-308, 1998.
- 503 Elbert, W., Taylor, P. E., Andreae, M. O., and Pöschl, U.: Contribution of fungi to primary biogenic aerosols in the
- 504 atmosphere: wet and dry discharged spores, carbohydrates, and inorganic ions, Atmos. Chem. Phys., 7, 4569-4588, 2007.
- 505 Estillore, A. D., Trueblood, J. V., and Grassian, V. H.: Atmospheric chemistry of bioaerosols: heterogeneous and 506 multiphase reactions with atmospheric oxidants and other trace gases, Chem. Sci., 7, 6604-6616, 2016.
- 507 Fierer, N., Jackson, J. A., Vilgalys, R., and Jackson, R. B.: Assessment of soil microbial community structure by use of 508 taxon-specific quantitative PCR assays, Appl. Environ. Microb., 71, 4117-4120, 2005.
- 509 Fröhlich-Nowoisky, J., Pickersgill, D. A., Després, V. R., and Pöschl, U.: High diversity of fungi in air particulate matter,
- 510 P. Natl. Acad. Sci. USA, 106, 12814-12819, 2009.

Atmos. Chem. Phys. Discuss., https://doi.org/10.5194/acp-2017-620 Manuscript under review for journal Atmos. Chem. Phys. Discussion started: 21 July 2017

© Author(s) 2017. CC BY 4.0 License.





- 511 Fröhlich-Nowoisky, J., Burrows, S. M., Xie, Z., Engling, G., Solomon, P. A., Fraser, M. P., Mayol-Bracero, O. L., Artaxo, 512 P., Begerow, D., Conrad, R., Andreae, M. O., Després, V. R., and Pöschl, U.: Biogeography in the air: fungal diversity
- 513 over land and oceans, Biogeosciences, 9, 1125-1136, 2012.
- Fröhlich-Nowoisky, J., Kampf, C. J., Weber, B., Huffman, J. A., Pöhlker, C., Andreae, M. O., Lang-Yona, N., Burrows, 514
- 515 S. M., Gunthe, S. S., and Elbert, W.: Bioaerosols in the Earth system: Climate, health, and ecosystem interactions, Atmos.
- 516 Res., 182, 346-376, 2016.
- 517 Georgakopoulos, D., Després, V., Fröhlich-Nowoisky, J., Psenner, R., Ariya, P., Pósfai, M., Ahern, H., Moffett, B., and
- 518 Hill, T.: Microbiology and atmospheric processes: biological, physical and chemical characterization of aerosol particles, 519 Biogeosciences Discussions, 5, 1469-1510, 2008.
- 520 Gosselin, M. I., Rathnayake, C. M., Crawford, I., Pöhlker, C., Fröhlich-Nowoisky, J., Schmer, B., Després, V. R., Engling, 521 G., Gallagher, M., and Stone, E.: Fluorescent bioaerosol particle, molecular tracer, and fungal spore concentrations during
- 522
- dry and rainy periods in a semi-arid forest, Atmos. Chem. Phys., 16, 15165-15184, 2016.
- 523 Haga, D., Burrows, S. M., Iannone, R., Wheeler, M., Mason, R., Chen, J., Polishchuk, E., Pöschl, U., and Bertram, A. K.:
- 524 Ice nucleation by fungal spores from the classes Agaricomycetes, Ustilaginomycetes, and Eurotiomycetes, and the effect 525 on the atmospheric transport of these spores, Atmos. Chem. Phys., 14, 8611-8630, 2014.
- 526 Hari, P., and Kulmala, M.: Station for Measuring Ecosystem-Atmosphere Relations (SMEAR II), Boreal Environ. Res.,
- 527 10, 315-322, 2005.
- 528 Harrel, F.E: Hmisc: Harrell miscellaneous, R package version 4.0-2, available at: https://CRAN.R-529 project.org/package=Hmisc (last access: 27 March 2017), 2016
- 530 Hassett, M. O., Fischer, M. W., and Money, N. P.: Mushrooms as rainmakers: how spores act as nuclei for raindrops, 531 PloS one, 10, e0140407, 2015.
- 532 Hirano, S. S., Baker, L. S., and Upper, C. D.: Raindrop momentum triggers growth of leaf-associated populations of 533 Pseudomonas syringae on field-grown snap bean plants, Appl. Environ. Microb., 62, 2560-2566, 1996.
- 534 Hock, N., Schneider, J., Borrmann, S., Römpp, A., Moortgat, G., Franze, T., Schauer, C., Pöschl, U., Plass-Dülmer, C., 535 and Berresheim, H.: Rural continental aerosol properties and processes observed during the Hohenpeissenberg Aerosol
- 536 Characterization Experiment (HAZE2002), Atmos. Chem. Phys., 8, 603-623, 2008.
- 537 Huffman, J. A., Prenni, A., DeMott, P., Pöhlker, C., Mason, R., Robinson, N., Fröhlich-Nowoisky, J., Tobo, Y., Després,
- 538 V., and Garcia, E.: High concentrations of biological aerosol particles and ice nuclei during and after rain, Atmos. Chem. 539 Phys., 13, 6151-6164, 2013.
- 540 Hussein, T., Norros, V., Hakala, J., Petäjä, T., Aalto, P. P., Rannik, Ü., Vesala, T., and Ovaskainen, O.: Species traits and 541 inertial deposition of fungal spores, J. Aerosol Sci., 61, 81-98, 2013.
- 542 Jaenicke, R.: Abundance of cellular material and proteins in the atmosphere, Science, 308, 73-73, 2005.
- 543 Jones, A. M., and Harrison, R. M.: The effects of meteorological factors on atmospheric bioaerosol concentrations-a 544 review, Sci. Total Environ., 326, 151-180, 2004.
- 545 Junninen, H., Lauri, A., Keronen, P., Aalto, P., Hiltunen, V., Hari, P., and Kulmala, M.: Smart-SMEAR: on-line data 546 exploration and visualization tool tor SMEAR stations, Boreal Environ. Res., 14, 447-457, 2009.
- 547 Kauserud, H., Heegaard, E., Halvorsen, R., Boddy, L., Høiland, K., and Stenseth, N. C.: Mushroom's spore size and time 548 of fruiting are strongly related: is moisture important?, Biology Lett., 7, 273-276, 2011.
- 549 Kuuskeri, J., Häkkinen, M., Laine, P., Smolander, O.-P., Tamene, F., Miettinen, S., Nousiainen, P., Kemell, M., Auvinen, 550 P., and Lundell, T.: Time-scale dynamics of proteome and transcriptome of the white-rot fungus Phlebia radiata: growth 551
- on spruce wood and decay effect on lignocellulose, Biotechnol. Biofuels, 9, 192, 2016.
- 552 Laakso, L., Hussein, T., Aarnio, P., Komppula, M., Hiltunen, V., Viisanen, Y., and Kulmala, M.: Diurnal and annual 553 characteristics of particle mass and number concentrations in urban, rural and Arctic environments in Finland, Atmos. 554 555 Environ., 37, 2629-2641, 2003.
- Lighthart, B.: Mini-review of the concentration variations found in the alfresco atmospheric bacterial populations, 556 Aerobiologia, 16, 7-16, 2000.
- 557 Madsen, A. M.: Effects of airflow and changing humidity on the aerosolization of respirable fungal fragments and conidia 558 of Botrytis cinerea, Appl. Environ. Microb., 78, 3999-4007, 2012.
- 559 Manninen, H. E., Bäck, J., Sihto-Nissilä, S.-L., Huffman, J. A., Pessi, A.-M., Hiltunen, V., Aalto, P. P., Hidalgo 560 Fernández, P. J., Hari, P., and Saarto, A.: Patterns in airborne pollen and other primary biological aerosol particles
- 561 (PBAP), and their contribution to aerosol mass and number in a boreal forest, Boreal Environ. Res., 19, 383-405, 2014.
- 562 Mason, R. H., Si, M., Chou, C., Irish, V., Dickie, R., Elizondo, P., Wong, R., Brintnell, M., Elsasser, M., and Lassar, W.:
- 563 Size-resolved measurements of ice-nucleating particles at six locations in North America and one in Europe, Atmos. 564 Chem. Phys., 16, 1637-1651, 2016.
- 565 Matos, J. T., Duarte, R. M., and Duarte, A. C.: Challenges in the identification and characterization of free amino acids
- 566 and proteinaceous compounds in atmospheric aerosols: A critical review, TrAC Trends Anal. Chem., 75, 97-107, 2016.
- 567 McGregor, K. G., and Anastasio, C.: Chemistry of fog waters in California's Central Valley: 2. Photochemical 568 transformations of amino acids and alkyl amines, Atmos. Environ., 35, 1091-1104, 2001.
- 569 Milne, P. J., and Zika, R. G.: Amino acid nitrogen in atmospheric aerosols: Occurrence, sources and photochemical
- 570 modification, J. Atmos. Chem., 16, 361-398, 1993.

Atmos. Chem. Phys. Discuss., https://doi.org/10.5194/acp-2017-620 Manuscript under review for journal Atmos. Chem. Phys. Discussion started: 21 July 2017

© Author(s) 2017. CC BY 4.0 License.





571 Möhler, O., DeMott, P. J., Vali, G., and Levin, Z.: Microbiology and atmospheric processes: the role of biological particles 572 in cloud physics, Biogeosciences, 4, 1059-1071, 2007.

- 573 Möhler, O., Georgakopoulos, D., Morris, C., Benz, S., Ebert, V., Hunsmann, S., Saathoff, H., Schnaiter, M., and Wagner,
- 574 R.: Heterogeneous ice nucleation activity of bacteria: new laboratory experiments at simulated cloud conditions,
- 575 Biogeosciences, 5, 1425-1435, 2008.
- 576 Morris, C., Sands, D., Bardin, M., Jaenicke, R., Vogel, B., Leyronas, C., Ariya, P., and Psenner, R.: Microbiology and
- 577 atmospheric processes: research challenges concerning the impact of airborne micro-organisms on the atmosphere and
- 578 climate, Biogeosciences, 8, 17-25, 2011.
- 579 Morris, C. E., Soubeyrand, S., Bigg, E. K., Creamean, J. M., and Sands, D. C.: Mapping rainfall feedback to reveal the 580 potential sensitivity of precipitation to biological aerosols, B. Am. Meteorol. Soc., 2016.
- 581 Mumford, R. A., Lipke, H., Laufer, D. A., and Feder, W. A.: Ozone-induced changes in corn pollen, Environ. Sci. 582 Technol., 6, 427-430, 1972.
- 583 Oksanen, J. F.; Blanchet, G.; Kindt, R.; Legendre, P.; Minchin, P. R.; O'Hara, R. B.; Simpson, G. L.; Solymos, P.; Henry,
- 584 M.; Stevens, M. H. H.; Wagner, H.: vegan: Community Ecology Package, R package version 2.3-3, available at: 585 https://CRAN.R-project.org/package=vegan (last access: 27 March 2017), 2016.
- 586 Ozler, H., Pehlivan, S., and Bayrak, F.: Analysis of free amino acid and total protein content in pollen of some allergenic 587 taxa, Asian J. Plant Sci., 8, 308, 2009.
- 588 Pietsch, R. B., David, R. F., Marr, L. C., Vinatzer, B., and Schmale III, D. G.: Aerosolization of Two Strains (Ice+ and 589
- Ice-) of Pseudomonas syringae in a Collison Nebulizer at Different Temperatures, Aerosol Sci. Tech., 49, 159-166, 2015. 590 Pöschl, U., Martin, S., Sinha, B., Chen, Q., Gunthe, S., Huffman, J., Borrmann, S., Farmer, D., Garland, R., and Helas,
- 591 G.: Rainforest aerosols as biogenic nuclei of clouds and precipitation in the Amazon, Science, 329, 1513-1516, 2010.
- 592 Prenni, A., Tobo, Y., Garcia, E., DeMott, P., Huffman, J., McCluskey, C., Kreidenweis, S., Prenni, J., Pöhlker, C., and
- 593 Pöschl, U.: The impact of rain on ice nuclei populations at a forested site in Colorado, Geophys. Res. Lett., 40, 227-231, 594 2013
- 595 Puc, M.: Characterisation of pollen allergens, Ann. Agr. Env. Med., 10, 143-150, 2003.
- 596 Pummer, B., Bauer, H., Bernardi, J., Bleicher, S., and Grothe, H.: Suspendable macromolecules are responsible for ice 597 nucleation activity of birch and conifer pollen, Atmos. Chem. Phys., 12, 2541-2550, 2012.
- 598 Pummer, B., Budke, C., Augustin-Bauditz, S., Niedermeier, D., Felgitsch, L., Kampf, C., Huber, R., Liedl, K., Loerting, 599 T., and Moschen, T.: Ice nucleation by water-soluble macromolecules, Atmos. Chem. Phys., 15, 4077-4091, 2015.
- 600 Purohit, H. J., Raje, D. V., and Kapley, A.: Identification of signature and primers specific to genus Pseudomonas using 601 mismatched patterns of 16S rDNA sequences, BMC bioinformatics, 4, 19, 2003.
- Rathnayake, C. M., Metwali, N., Jayarathne, T., Kettler, J., Huang, Y., Thorne, P. S., O'Shaughnessy, P. T., and Stone, 602 603 E. A.: Influence of rain on the abundance of bioaerosols in fine and coarse particles, Atmos. Chem. Phys., 17, 2459-2475,
- 604 2017.
- 605 R Core Team: R: A language and environment for statistical computing, R Foundation for Statistical Computing, Vienna,
- 606 Austria, available at: https://www.r-project.org/ (last access: 17 June 2017), 2016.
- 607 Reponen, T., Willeke, K., Ulevicius, V., Reponen, A., and Grinshpun, S. A.: Effect of relative humidity on the 608 aerodynamic diameter and respiratory deposition of fungal spores, Atmos. Environ., 30, 3967-3974, 1996.
- 609 Reponen, T., Grinshpun, S., Conwell, K., Wiest, J., and Anderson, M.: Aerodynamic versus physical size of spores: 610 measurement and implication for respiratory deposition, Grana, 40, 119-125, 2001.
- 611 Rolph, G.: Real-time Environmental Applications and Display sYstem (READY)
- Website 612 (http://www.arl.noaa.gov/ready/hysplit4. html), NOAA Air Resources Laboratory, Silver Spring, Md, 2003.
- 613 Saarnio, K., Niemi, J. V., Saarikoski, S., Aurela, M., Timonen, H., Teinila, K., Myllynen, M., Freyi, A., Lamberg, H.,
- 614 and Jokiniemi, J.: Using monosaccharide anhydrides to estimate the impact of wood combustion on fine particles in the
- 615 Helsinki Metropolitan Area, Boreal Environ. Res., 17, 163-184, 2012.
- 616 Samy, S., Robinson, J., and Hays, M. D.: An advanced LC-MS (Q-TOF) technique for the detection of amino acids in 617 atmospheric aerosols, Anal. Bioanal. Chem., 401, 3103-3113, 2011.
- 618 Samy, S., Robinson, J., Rumsey, I. C., Walker, J. T., and Hays, M. D.: Speciation and trends of organic nitrogen in 619 southeastern US fine particulate matter (PM2.5), J. Geophys. Res. Atmos., 118, 1996-2006, 2013.
- 620 Scalabrin, E., Zangrando, R., Barbaro, E., Kehrwald, N., Gabrieli, J., Barbante, C., and Gambaro, A.: Amino acids in 621 Arctic aerosols, Atmos. Chem. Phys., 12, 10453-10463, 2012.
- 622 Schneider, J., Freutel, F., Zorn, S., Chen, Q., Farmer, D., Jimenez, J., Martin, S., Artaxo, P., Wiedensohler, A., and 623 Borrmann, S.: Mass-spectrometric identification of primary biological particle markers and application to pristine 624 submicron aerosol measurements in Amazonia, Atmos. Chem. Phys., 11, 11415-11429, 2011.
- 625 Schumacher, C. J., Pöhlker, C., Aalto, P., Hiltunen, V., Petäjä, T., Kulmala, M., Pöschl, U., and Huffman, J. A.: Seasonal
- 626 cycles of fluorescent biological aerosol particles in boreal and semi-arid forests of Finland and Colorado, Atmos. Chem.
- 627 Phys., 13, 11987-12001, 2013.
- 628 Sesartic, A., Lohmann, U., and Storelymo, T.: Bacteria in the ECHAM5-HAM global climate model, Atmos. Chem.
- 629 Phys., 12, 8645-8661, 2012.

Atmos. Chem. Phys. Discuss., https://doi.org/10.5194/acp-2017-620 Manuscript under review for journal Atmos. Chem. Phys. Discussion started: 21 July 2017







630 Sesartic, A., Lohmann, U., and Storelymo, T.: Modelling the impact of fungal spore ice nuclei on clouds and precipitation, 631 Environ. Res. Lett., 8, 014029, 2013.

- 632 Spracklen, D., and Heald, C.: The contribution of fungal spores and bacteria to regional and global aerosol number and
- 633 ice nucleation immersion freezing rates, Atmos. Chem. Phys., 14, 9051-9059, 2014.
- 634 Staton, S. J., Woodward, A., Castillo, J. A., Swing, K., and Hayes, M. A.: Ground level environmental protein
- 635 concentrations in various ecuadorian environments: Potential uses of aerosolized protein for ecological research, Ecol. Indic., 48, 389-395, 2015. 636
- 637 Stein, A. F., Draxler, R. R., Rolph, G. D., Stunder, B. J. B., Cohen, M. D., and Ngan, F.: NOAA's HYSPLIT Atmospheric 638 Transport and Dispersion Modeling System, B. Am. Meteorol. Soc., 96, 2059-2077, 2015.
- 639 Stoddard, S. F., Smith, B. J., Hein, R., Roller, B. R., and Schmidt, T. M.: rrnDB: improved tools for interpreting rRNA 640 gene abundance in bacteria and archaea and a new foundation for future development, Nucleic Acids Res., 43, D593-641 D598, 2014.
- 642 Taylor, P., Flagan, R., Miguel, A., Valenta, R., and Glovsky, M.: Birch pollen rupture and the release of aerosols of 643 respirable allergens, Clin. Exp. Allergy, 34, 1591-1596, 2004.
- 644 Vainio, E. J., and Hantula, J.: Direct analysis of wood-inhabiting fungi using denaturing gradient gel electrophoresis of 645 amplified ribosomal DNA, Mycol. Res., 104, 927-936, 2000.
- 646 Violaki, K., and Mihalopoulos, N.: Water-soluble organic nitrogen (WSON) in size-segregated atmospheric particles over 647 the Eastern Mediterranean, Atmos. Environ., 44, 4339-4345, 2010.
- 648 Visez, N., Chassard, G., Azarkan, N., Naas, O., Sénéchal, H., Sutra, J.-P., Poncet, P., and Choël, M.: Wind-induced 649 mechanical rupture of birch pollen: potential implications for allergen dispersal, J. Aerosol Sci., 89, 77-84, 2015.
- 650 Wei, T.; Simko, V.: corrplot: Visualization of a Correlation Matrix, R package version 0.77, available at: https://CRAN.R-651 project.org/package=corrplot (last access: 27 March 2017), 2016
- 652 Zangrando, R., Barbaro, E., Kirchgeorg, T., Vecchiato, M., Scalabrin, E., Radaelli, M., Đorđević, D., Barbante, C., and
- 653 Gambaro, A.: Five primary sources of organic aerosols in the urban atmosphere of Belgrade (Serbia), Sci. Total Environ., 654 571, 1441-1453, 2016.
- 655 Zhang, Q., Anastasio, C., and Jimenez-Cruz, M.: Water-soluble organic nitrogen in atmospheric fine particles (PM2. 5) 656 from northern California, J. Geophys. Res. Atmos., 107 (D11), 2002.
- 657 Zhang, Q., and Anastasio, C.: Free and combined amino compounds in atmospheric fine particles (PM 2.5) and fog waters 658 from Northern California, Atmos. Environ., 37, 2247-2258, 2003.
- 659 Zhang, T., Engling, G., Chan, C.-Y., Zhang, Y.-N., Zhang, Z.-S., Lin, M., Sang, X.-F., Li, Y., and Li, Y.-S.: Contribution 660 of fungal spores to particulate matter in a tropical rainforest, Environ. Res. Lett., 5, 024010, 2010.
- 661 Zhu, C., Kawamura, K., and Kunwar, B.: Organic tracers of primary biological aerosol particles at subtropical Okinawa 662 Island in the western North Pacific Rim, J. Geophys. Res. Atmos., 120, 5504-5523, 2015.
- 663 664

665

666 Table 1. Average (±standard deviation) concentration and range of each component measured and the number of filter

667 samples (n) analyzed in total.

| Component                           | <1 µm   | 1-2.5 μm   | 2.5-10 µm  | $>10 \ \mu m$  | total particles    |
|-------------------------------------|---|--|--|--|--------------------|
| PM (μg/m <sup>3</sup> )             | 5.9±4.1   | 2.1±2.3  | 2.4±3.2  | 2.1±3.5  | 12.3±12.2          |
| n=138                               | (1.0-18.1)  | (0.2-11.9)   | (0.2-15.9)   | (0.03-17.4)  | (2.2-63.3)         |
| FAA (ng/m <sup>3</sup> )            | $5.22 \pm 10.11$  | $10.95 \pm 31.54$  | $18.45 \pm 53.55$  | $27.62 \pm 85.71$  | $57.91{\pm}174.17$ |
| n=69                                | ( <loq-41.21)< td=""><td>(<loq-120.24)< td=""><td>(<loq-224.92)< td=""><td>(0.04-364.65)</td><td>(1.13-751.01)</td></loq-224.92)<></td></loq-120.24)<></td></loq-41.21)<> | ( <loq-120.24)< td=""><td>(<loq-224.92)< td=""><td>(0.04-364.65)</td><td>(1.13-751.01)</td></loq-224.92)<></td></loq-120.24)<> | ( <loq-224.92)< td=""><td>(0.04-364.65)</td><td>(1.13-751.01)</td></loq-224.92)<>  | (0.04-364.65)  | (1.13-751.01)      |
| DNA (ng/m <sup>3</sup> )            |   | $5.16\pm5.20$  | $1.99 \pm 3.47$  | $2.18 \pm 3.76$  | $8.60{\pm}11.41$   |
| n=51                                | -   | ( <loq-18.56)< td=""><td>(0.002-14.67)</td><td>(<loq-14.80)< td=""><td>(0.24-48.04)</td></loq-14.80)<></td></loq-18.56)<>      | (0.002-14.67)  | ( <loq-14.80)< td=""><td>(0.24-48.04)</td></loq-14.80)<> | (0.24-48.04)       |
| Bacteria (cells/m <sup>3</sup> )    |   | 2811±6619  | 2171±4032  | 1341±3424  | 6323±13748         |
| n=51                                | -   | (17-27551)   | (56-16746)   | (41-14434)   | (137-58731)        |
| Pseudomonas (cells/m <sup>3</sup> ) |   | 74±125   | 4±5  | 13±26  | 86±122             |
| n=51                                | -   | (0.1-469)  | ( <loq-14)< td=""><td>(<loq-84)< td=""><td>(1-469)</td></loq-84)<></td></loq-14)<> | ( <loq-84)< td=""><td>(1-469)</td></loq-84)<>            | (1-469)            |
| Fungi (cells/m <sup>3</sup> )       |   | 4022±7518  | 5579±10614   | 648±1809   | $10173 \pm 15622$  |
| n=51                                | -   | (2-27838)  | (9-42667)  | ( <loq-7129)< td=""><td>(27-55839)</td></loq-7129)<>     | (27-55839)         |