

Characterization of organic nitrate constituents of secondary organic aerosol (SOA) from nitrate radical-initiated oxidation of limonene using High-Resolution Chemical Ionization Mass Spectrometry

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Abstract: The gas phase nitrate radical (NO_3^\bullet) initiated oxidation of limonene can produce organic nitrate species with varying physical properties. Low-volatility products can contribute to secondary organic aerosol (SOA) formation and organic nitrates may serve as a NO_x reservoir, which could be especially important in regions with high biogenic emissions. This work presents the measurement results from flow reactor studies on the reaction of NO_3^\bullet with limonene using a High-Resolution Time-of-Flight Chemical Ionization Mass Spectrometer (HR-ToF-CIMS) combined with a Filter Inlet for Gases and AEROSols (FIGAERO). Major condensed-phase species were compared to those in the Master Chemical Mechanism (MCM) limonene mechanism, and many non-listed species were identified. The volatility properties of the most prevalent organic nitrates in the produced SOA were determined. Analysis of multiple experiments resulted in the identification of several dominant species (including $\text{C}_{10}\text{H}_{15}\text{NO}_6$, $\text{C}_{10}\text{H}_{17}\text{NO}_6$, $\text{C}_8\text{H}_{11}\text{NO}_6$, $\text{C}_{10}\text{H}_{17}\text{NO}_7$, and $\text{C}_9\text{H}_{13}\text{NO}_7$) that occurred in the SOA under all conditions considered. Additionally, the formation of dimers was consistently observed and these species resided almost completely in the particle phase. The identities of these species are discussed, and formation mechanisms are proposed. Cluster analysis of the desorption temperatures corresponding to the analyzed particle-phase species yielded at least five distinct groupings based on a combination of molecular weight and desorption profile. Overall, the results indicate that the oxidation of limonene by NO_3^\bullet produces a complex mixture of highly oxygenated monomer and dimer products that contribute to SOA formation.

1 Introduction

Oxidation of gas-phase organic species contribute significantly to particle formation and growth (Hallquist et al., 2009; Smith et al., 2008; Wehner et al., 2005), and thus a thorough understanding of secondary organic aerosol (SOA) formation mechanisms is important for the accurate estimation of its impact on the climate system (Kanakidou et al., 2005).

Secondary organic aerosols form primarily via the photooxidation of volatile organic compounds (VOCs), yielding less volatile products, which can then partition into the condensed phase (Hallquist et al., 2009; Kroll and Seinfeld, 2008), especially when pre-existing aerosols (e.g., inorganic seed particles) are present (Kroll et al., 2007). The products resulting from atmospheric oxidation may be classified as low volatility, semi-volatile, and intermediate volatility OCs, i.e., LVOCs, SVOCs, and IVOCs, respectively (Donahue et al., 2012; Jimenez et al., 2006; Murphy et

37 al., 2014). In addition, extremely low volatility OCs (i.e., ELVOCs) contribute significantly to
38 aerosol formation and early growth (Ehn et al., 2014; Jokinen et al., 2015). The oxidation of
39 VOCs by the primary atmospheric oxidants, O₃ and [•]OH, has been extensively investigated (Cao
40 and Jang, 2008; Hallquist et al., 2009; Kanakidou et al., 2005; Kroll and Seinfeld, 2008).
41 Although less studied than the photo-oxidation of VOCs, the reaction of VOCs with the nitrate
42 radical (NO₃[•]) and the resulting formation of organic nitrates are also important, especially for
43 nocturnal chemistry (Roberts, 1990, Brown and Stutz, 2012; Perring et al., 2013; Kiendler-
44 Scharr et al., 2016, Ng et al., 2017). Significant concentrations of these nitrates have been
45 detected in the gas and condensed phases in both field and laboratory studies (Ayres et al., 2015;
46 Beaver et al., 2012; Boyd et al., 2017; Bruns et al., 2010; Day et al., 2010, Fry et al., 2013; Lee et
47 al., 2016; Nah et al., 2016; Paulot et al., 2009; Rindelaub et al., 2014, 2015, Rollins et al., 2012,
48 2013; Xu et al., 2016, Kiendler-Scharr et al., 2016).

49 Organic nitrates (RONO₂) and organic peroxy nitrates (RO₂NO₂), such as peroxy acetyl
50 nitrate (PAN), may also form in the atmosphere (Roberts, 1990; Singh and Hanst, 1981; Temple
51 and Taylor, 1983). RO₂NO₂ may form via the reaction of organic peroxy nitrates (RO₂[•]) with
52 NO₂, while RONO₂ may form directly through either the reaction of RO₂[•] with NO or the
53 reaction of unsaturated VOCs with NO₃[•].

54 Secondary organic aerosol-precursor VOCs arise mainly from the emission and reaction of
55 biogenic VOCs (BVOCs) (Hallquist et al., 2009), with up to 90% of the global VOC budget
56 originating from biogenic sources (Glasius and Goldstein, 2016; Guenther et al., 1995). Isoprene,
57 the main constituent of global BVOC terrestrial emissions (600 Tg yr⁻¹) (Guenther et al., 2006), is
58 highly reactive with [•]OH, O₃, and NO₃[•] (Atkinson et al., 1995; Hallquist et al., 2009). However,
59 monoterpenes typically have higher SOA yields than isoprene (Carlton et al., 2009; Presto et al.,
60 2005b) and regarding atmospheric emissions, α-pinene, β-pinene, and limonene constitute the
61 main monoterpenes emitted into the atmosphere (Guenther et al., 2012). In addition to its high
62 emission rates, limonene is especially interesting as a model BVOC, due to its relatively high
63 reaction rates (Ziemann and Atkinson, 2012) and occurrence in indoor environments, owing to
64 emission sources, such as air fresheners and other household products (Wainman et al., 2000).

65 The reactions and mechanisms of α-pinene and β-pinene oxidation have been more
66 thoroughly studied (Bonn and Moorgat, 2002; Presto et al., 2005a, 2005b; Fry et al., 2009;
67 Perraud et al., 2010) than those associated with limonene. Several studies have focused on the
68 ozonolysis of and SOA formation from limonene (Leungsakul et al., 2005; Jonsson et al., 2006,
69 2008a; Zhang et al., 2006; Baptista et al., 2011; Sun et al., 2011; Pathak et al., 2012; Jiang et al.,
70 2013; Youssefi and Waring, 2014;). NO₃[•] oxidation of limonene and the resulting organic nitrates
71 that may contribute to SOA formation have, however, rarely been investigated (Hallquist et al.,
72 1999; Spittler et al., 2006 Fry et al., 2011, 2014; Boyd et al., 2017). In relation to the reaction with
73 NO₃[•], major non-nitrate products of limonene (including endolim) have been identified, but
74 significant SOA formation was preceded by the occurrence of multiple unidentified nitrates
75 (Hallquist et al., 1999; Spittler et al., 2006). Moreover, although mechanistic models and

76 molecular identities of these products have been proposed, direct measurement and identification
77 thereof have yet to be reported. Further elucidation of the mechanisms governing and products
78 generated by the reactions of limonene and NO_3^\bullet are warranted, since organic nitrates from
79 BVOCs (including limonene) have been consistently observed in field studies (Perring et al.,
80 2009; Ayres et al., 2015; Beaver et al., 2012; Lee et al., 2016, 2014b;).

81 Additionally, the contribution of low-volatility products to the SOA mass may increase
82 with the formation of dimers from aerosol components generated by VOC oxidation. Numerous
83 dimers or oligomers have been found in SOA generated by monoterpene species (e.g.
84 Emanuelsson et al., 2013; Kourtchev et al., 2014, 2016; Kristensen et al., 2016; Müller et al.,
85 2007; Tolocka et al., 2004). However, the speciation of observed dimers and oligomers from
86 organic nitrates, especially with respect to detailed formation mechanisms, has rarely been
87 reported.

88 Here we report the chemical composition of low-volatility gas and aerosol-phase species,
89 formed from mixtures of N_2O_5 and limonene, as measured by a High Resolution Time-of-Flight
90 Chemical Ionization Mass Spectrometer (HR-ToF-CIMS) coupled to a Filter Inlet for Gases and
91 AEROSols (FIGAERO) inlet (Lopez-Hilfiker et al., 2014). The objectives of this work were
92 three-fold namely, to: (i) determine the molecular formulae of major nitrate species produced
93 from the reaction of limonene with NO_3^\bullet , that could contribute significantly to SOA formation
94 and growth, (ii) compare the distribution of measured products to that of the expected products
95 (based on the Master Chemical Mechanism (MCM)) to identify any discrepancies in the
96 mechanistic understanding of nitrate formation from limonene, and (iii) categorize, via cluster
97 analysis, the thermodynamic desorption data measured for selected condensed-phase species.

98 **2 Methods**

99 **2.1 Experimental setup**

100 Experiments were performed in the Gothenburg Flow Reactor for Oxidation Studies at
101 low Temperatures (GFROST) at the University of Gothenburg. In previous studies, this facility
102 was used for studying the impact of relative humidity, OH-scavengers, and temperature on SOA
103 formation via monoterpene ozonolysis (Emanuelsson et al., 2013; Jonsson et al., 2008a, 2008b),
104 its volatility properties (Pathak et al., 2012), and dimer formation during the ozonolysis of α -
105 pinene (Kristensen et al., 2016). The inflow of zero air and the reagents is fixed at a total flow of
106 1.6 L per min (LPM). The experiments are all run at low RH ($\leq 1\%$) and a constant temperature of
107 20°C . To catch only the center portion of the laminar flow and avoid unnecessary interference
108 from wall effects, samples are taken through a cone at the end of the reactor at 0.95 LPM. The
109 average residence time of the sampled portion of the mixture is 240 s. Due to the flow
110 restrictions, a make-up flow of zero air is added to the sample, immediately after the outlet, prior
111 to being sampled by the instruments. The amount of dilution flow necessary is constrained by the
112 flow required by the HR-ToF-CIMS. Figure 1 shows a diagram of the experimental setup.

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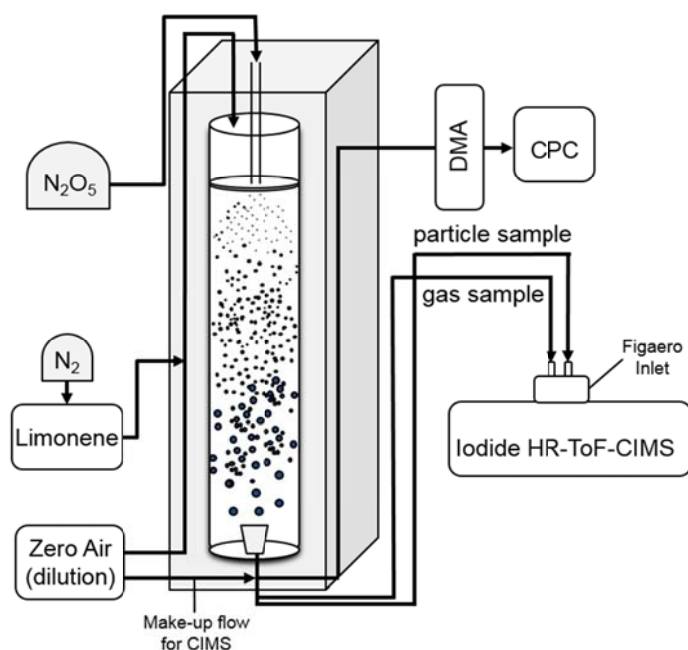


Figure 1. Diagram of experimental setup of GFROST during experiments.

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116 Gas and particle-phase products were measured using a High-Resolution Time-of-Flight
 117 Chemical Ionization Mass Spectrometer (HR-ToF-CIMS) coupled to a Filter Inlet for Gases and
 118 AEROSols (FIGAERO) (Lopez-Hilfiker et al., 2014). The HR-ToF-CIMS can be operated in
 119 either negative- or positive-ionization modes, using various reagent-ion sources. CIMS
 120 measurement techniques have previously been employed for the measurement of organic nitrate
 121 products of monoterpenes (Beaver et al., 2012; Paulot et al., 2009) using multiple reagent ions
 122 (Lee et al., 2014a). In this work, the HR-ToF-CIMS was operated using negative Iodide (I^-) ion
 123 as the reagent in all experiments. Dry UHP N_2 was passed over a permeation tube containing
 124 liquid CH_3I (Alfa Aesar, 99%), and $I(H_2O)_n^-$ ions were generated by directing the flow over a
 125 ^{210}Po radioactive source. Reaction products (e.g., species X) were identified by their
 126 corresponding cluster ions, XI^- , thereby allowing the collection of whole-molecule data. The
 127 reagent and sample flowed into the Ion-Molecule Reaction (IMR) chamber of the instrument at a
 128 nominal individual rate of 2 LPM. The IMR was temperature-controlled at $40^\circ C$ and operated at
 129 a nominal pressure of 200 mbar. With I^- ionization, the sensitivity of a detected species (i.e., hz
 130 ppt^{-1}) can vary significantly with relative humidity (Lee et al., 2014a). However, the experiments
 131 were all performed at low RH ($\leq 1\%$) and, hence, the same sensitivity was realized for all the
 132 conditions considered.

133 The FIGAERO inlet was used during the experiments, and particles were collected on a
 134 Zefluor® PTFE membrane filter. The aerosol sample line and gas sample line were composed of
 135 12 mm copper tubing and 12 mm Teflon tubing, respectively. The inlet was operated in regular
 136 cycles – 1 h of gas-phase sampling and simultaneous particle collection, followed by a 1-h period
 137 where the filter was shifted into position over the IMR inlet and the collected SOA was desorbed.
 138 Desorption was facilitated by a 2 LPM flow of heated UHP N_2 over the filter. The temperature of

139 the N₂ was increased from 20 to 200°C in 50 min (3.5°C min⁻¹), and a subsequent 10-minute
140 temperature soak was performed to ensure complete removal of the remaining organic material
141 that volatilizes at 200°C. The measured species were distinguished based on their thermal
142 properties via the resulting desorption time-series profiles, hereafter referred to as thermograms.
143 Temperature gradients of >3.5°C min⁻¹ have been used in previous studies, but, in this work, a
144 lower gradient was used to enable optimum thermal separation (Lee et al., 2014a; Lopez-Hilfiker
145 et al., 2014). The HR-ToF-CIMS was configured to measure singly charged ions with a mass-to-
146 charge ratio (*m/z* or *Th*) of 7–720. Particles were contemporaneously sampled directly at the
147 outlet of the flow reactor, through a ¼" stainless steel 1 m sample line, by a Scanning Mobility
148 Particle Sizer (SMPS). The SMPS measured the number-size distribution used for estimating the
149 mass concentrations, based on the assumption of spherical particles with a density of 1.4 g cm⁻³
150 (Hallquist et al., 2009). In all cases, SOA was generated via nucleation and growth rather than by
151 using seed particles.

152 N₂O₅ was synthesized by reacting ≥20 ppm O₃ with pure NO₂ (98%, AGA Gas) in a glass
153 vessel and then passing the flow through a cold trap maintained at -78.5°C using dry ice. Even if
154 neither HNO₃ nor NO₂ was measured it is known from previous work that this method typically
155 provides a source with impurities less than a few percent. It is well known that the resulting white
156 solid would show signs of yellowing, due to nitric or nitrous acid contamination, if exposed to
157 moisture (e.g., ambient lab air) so handling of the N₂O₅ was done accordingly. The solid N₂O₅
158 was transferred to a diffusion vial fitted with a capillary tube (inner diameter: 2 mm). The N₂O₅
159 diffusion source was held at a constant temperature (-23 °C), and the gravimetrically determined
160 mass loss rate remained steady (r² value: 0.97–0.98) for several weeks. A similarly characterized
161 d-limonene (Alfa Aesar, 97%) diffusion source was held at temperatures ranging from 8.5 to
162 31.5°C and, using Gas Chromatography–Mass Spectrometry (GC-MS; Finnigan/Tremetrics),
163 diluted flow-reactor concentrations (15, 45, 92, and 150 ppb).

164 Experiments were performed over a range (1.0–113) of N₂O₅/limonene ratios (see Table 1
165 for a summary of experimental conditions). At a ratio around 1.0 one expects only the endocyclic
166 double bond to be reacting with NO₃ radicals while at higher ratio there is an increased
167 possibility for secondary chemistry where products will be susceptible for reaction with the NO₃
168 radical. For each set of conditions in the flow reactor, sampling was performed over a period of
169 6–12 h to ensure stability of conditions (e.g., gas-phase signals, total SOA mass) and repeatability
170 of the FIGAERO thermal-desorption cycles. An example of three sequential desorptions is shown
171 in Fig. S2.

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173 **Table 1.** Experimental conditions , concentrations and ratios of initial reactants and the resulting SOA mass.

#	N ₂ O ₅ (ppb)	Limonene (ppb)	N ₂ O ₅ / Limonene	Average SOA Mass* (µg m ³)
3	95	15	6.3	12±2
4	95	15	6.3	8±1
2	95	40	2.4	8±1
5	95	40	2.4	10±1
6	95	95	1	12±1
1	160	15	10.7	8±1
11	850	95	8.9	25±2
12	850	150	5.7	47±2
7	1700	15	113.3	7±1
8	1700	40	42.5	11±1
9	1700	95	17.9	43±2
10	1700	150	11.3	95±3

*Errors are given as standard deviation of the measured mean.

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177 2.2 CIMS data-analysis methods

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Data obtained from the HR-ToF-CIMS was analyzed using the Tofware (Tofwerk/Aerodyne) analysis software written in Igor Pro (WaveMetrics). High-resolution analysis allowed for ion identification with a resolution of ~4000 (m/Δm). Identified species were cross-checked with predicted species generated via the MCM v3.3.1 limonene mechanism (Saunders et al., 2003) and the corresponding theoretical product distribution was compared with the measured distribution for both gas and particle phase. For several ions, product formulas in the MCM were used as the major parameter for ion identification at a given *m/z*. However, this identification scheme resulted in the misidentification of several ions. The identification of high-mass ions (*m/z* > 500) was complicated by the fact that the number of possible formulas increases rapidly with increasing mass and carbon number of the ions. Nevertheless, the high accuracy of fits (≤5 ppm), where the identities of expected product ions were corroborated by the fits of expected isotopes, reduced uncertainties stemming from the mass calibration and provided reliable ion identifications. To further ensure the accuracy of the identities of high-mass ions, the fits of the identified ions were compared over all experiments.

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The high-resolution ion data was further analyzed with Python 3.5.2 using the pandas (McKinney, 2010, 2011) and NumPy (Van Der Walt et al., 2011) packages, and peaks in the ion thermograms were identified using an implementation of the PeakUtils package (v1.0.3, <http://pythonhosted.org/PeakUtils/>). For each experiment, the temperature (T_{max}) corresponding to the peak signal of each ion observed during the desorption of SOA particles was identified. Furthermore, a secondary temperature (T_{max,2}) was identified when double-peak behavior was observed.

199 2.3 Cluster-analysis methods

200 Cluster analysis, performed via the K-Means algorithm (scikit-learn machine learning
201 package; Pedregosa et al., 2011), was used to distinguish, based on their elemental composition
202 and thermodynamic behavior (T_{\max}), groups of ions observed during SOA desorption. This
203 algorithm, utilizing a random seeding approach (Arthur and Vassilvitskii, 2007), was chosen due
204 to the superior cluster separation realized after comparing several algorithms, including affinity
205 propagation and mean-shift clustering. The solution of the K-Means algorithm is obtained
206 through the minimization of an inertia function (see Eq. 1) Φ , which is equivalent to the sum of
207 the mean-squared distance between all samples and their corresponding cluster centroid, c
208 (Arthur and Vassilvitskii, 2007; Raschka, 2016). Here, $x^{(i)}$: sample (e.g., carbon number, oxygen
209 number, T_{\max}) in a set of n samples, $c^{(j)}$: cluster center of cluster j in a set of k clusters, and $w^{(i,j)}$:
210 weighting coefficient ($w^{(i,j)} = 1$ if $x^{(i)}$ is in cluster j , $w^{(i,j)}=0$ otherwise).

$$211 \quad \phi = \sum_{i=1}^n \sum_{j=1}^k w^{(i,j)} \|x^{(i)} - c^{(j)}\|^2 \quad (1)$$

212 The quality of the cluster separation was assessed through a silhouette score, $s(i)$
213 (Rousseeuw, 1987), which allows comparison of the intra-cluster and inter-cluster distances and,
214 for a sample i , is determined from:

$$215 \quad s(i) = \frac{b(i)-a(i)}{\max\{a(i),b(i)\}} \quad (2)$$

216 where, $a(i)$: average distance, or dissimilarity, between point i and each point within its
217 own cluster and $b(i)$: average dissimilarity between point i and all points within the nearest
218 neighboring cluster. The value of $s(i)$ ranges from -1 to 1 and reflects the quality of the clustering
219 with respect to the separation between members of each cluster. For example, a score of ~ 1
220 indicates that the point is relatively far away from the nearest neighboring cluster, while a score
221 of 0 suggests that the cluster separation is roughly equivalent to that of cohesion clusters; that is,
222 $a(i) \approx b(i)$. For all points within a clustered dataset, an average silhouette score can indicate the
223 adequacy of the cluster separation for a given number of clusters.

224 Detected ions were clustered based on their molecular weight (MW), elemental numbers
225 (n_C , n_H , n_O , n_N), and T_{\max} values. Compared with the other variables, MW and the carbon number
226 exhibited the highest correlation with T_{\max} . Clustering the ions based on these three variables
227 yielded the best separation with respect to mass and T_{\max} of the ions. Input variables were scaled
228 to values between 0 and 1 (based on their respective range of input values) to prevent any bias
229 associated with the relative magnitude of each variable (e.g., $MW \gg n_C$).

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231 **3 Results and discussion**

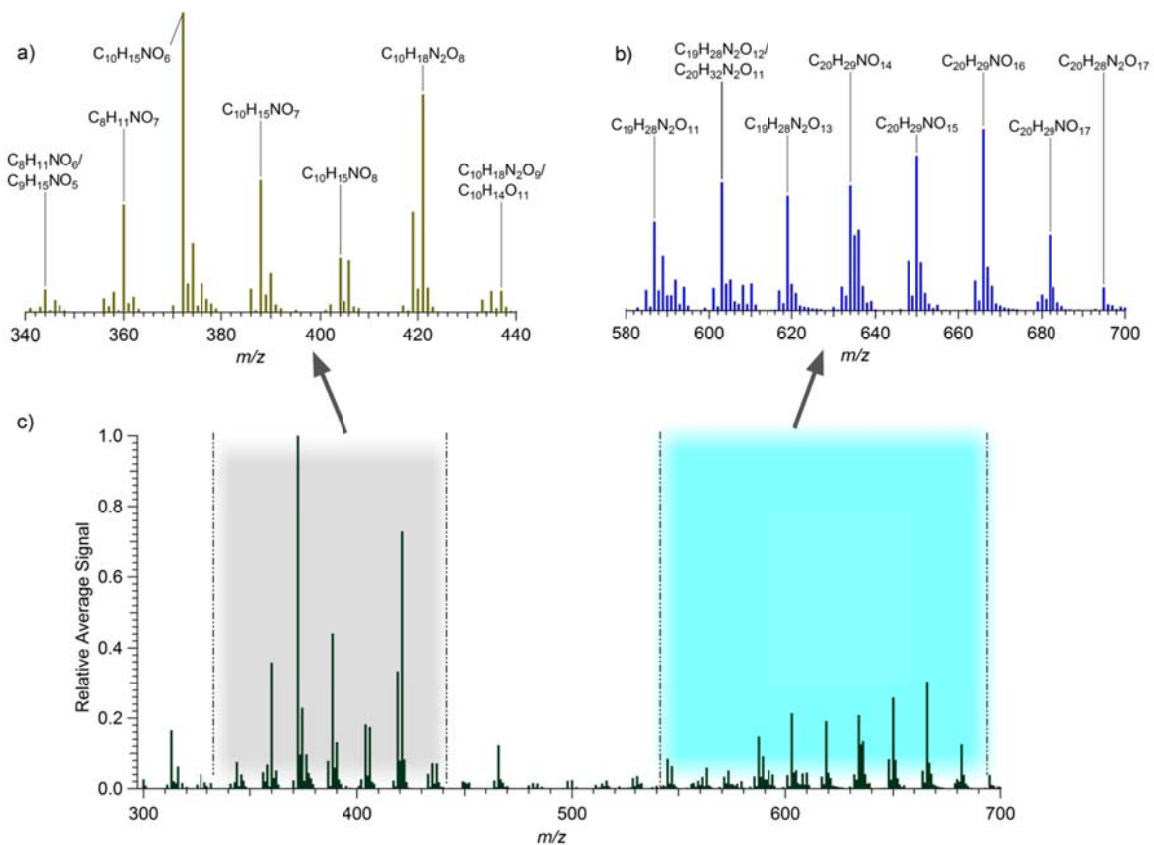
232 **3.1 Characterization of mass spectra from SOA and identification of species**

233 Products in both the gas and condensed phases were identified by analyzing HR-ToF-
234 CIMS data collected under various experimental conditions (Table 1). In each sampling regime,
235 major products were readily identifiable, and only modest or negligible fragmentation occurred
236 with application of the soft ionization technique. The focus in the current work was on condensed
237 phase products using the FIGARO inlet desorption. Recently, Stark *et al.* (2017) showed that
238 fragmentation during the desorption can occur within the FIGAERO. In the current work the
239 fragmentation within the FIGARO was not specifically investigated. However, from our cluster
240 analysis it was evidently that fragmentation occurred with specific features in e.g. molecular
241 weight and evaporation temperature. The ramp rate during desorption was therefore maintained
242 for all experiments to ensure, if fragmentation did occur, it would be consistent and enable
243 comparable analysis of the dataset. The mass-to-charge (m/z or Th) values of the most prominent
244 ions of species detected in the collected aerosol were determined from the average mass spectra
245 obtained during desorption cycles. The results revealed two distinct regions consisting of several
246 clusters of elevated ion signals (Fig. 2). These regions were present in all experiments (Table 1).
247 The occurrence of ions in these regions indicates a prevalence of lower-mass monomer species
248 (typically in the range m/z 340-440) and higher-mass dimer species (typically in the range m/z
249 580-700). These results are analogous to those of previous ozonolysis studies, where highly
250 oxygenated multifunctional (HOM) molecules from monoterpene oxidation were observed using
251 a nitrate HR-ToF-CIMS (Ehn *et al.*, 2014; Jokinen *et al.*, 2015; Mentel *et al.*, 2015). Figure 2
252 shows an average mass spectrum corresponding to four sequential 1-h desorption cycles of 12- μg
253 m^{-3} SOA samples from a reaction mixture with a N_2O_5 to limonene ratio of 2.4. The gas to
254 particle ratio of most ions were below one as illustrated in Fig. S1, whereas the focus of this work
255 was to characterize the particle phase.

256 In total, 198 of the identified organic ions constituted significant fractions of the aerosol
257 samples, but most of the signal emanated from only ~25% of these species. The dominant species
258 were identified by averaging the desorption-time series of all experiments and extracting the top
259 75th percentile (by averaging the signal during desorption) of the monomer and dimer ions. The
260 resulting set of ions consisted of 52 molecular species that accounted for 76% of the organic
261 signal during desorption, while the top 90th percentile of ions (20 ions) accounted for 56%. This
262 52-ion set consisted of 28 monomers ($C = 7-10$) and 24 dimers or oligomers ($C = 11-20$). From
263 the HR analysis the definition of monomer and dimer was specifically defined based on number
264 of carbons rather than the less strict used of the two m/z regions illustrated in Fig. 2. On average,
265 the top 75th percentile of monomers and the top 75th percentile of dimers accounted for 83% of
266 the total monomer signal and 70% of the total dimer signal, respectively. A full list of ions and
267 the composition of the 40th, 75th, and 90th percentile subsets can be found in the Supplementary
268 Information (Table S1). This list is based on a common sensitivity for detection that might not
269 always be true and highly variable (see e.g. Isaacman-Van Wertz *et al.*, 2017). However, with this
270 assumption the list will provide molecular identity of the most prominent organic compounds

271 contributing to the SOA mass outlined in Table 1. One could assess the contribution of these
272 peaks to the total mass loading, although with high variation in molecular mass and oxidation, the
273 sensitivity is likely to vary significantly, resulting in large error margins and therefore deeming
274 any interpretation highly speculative.

275 The lower-mass region, of the two mass-spectra regions (see Fig. 2) typically occurred at m/z
276 values ranging from 340 to 440 and containing mainly monomers. Several ions in this region
277 matched the predicted molecular formulas associated with the MCM limonene mechanism, and
278 the largest signals occurred for species consisting of 8–10 carbon atoms. E.g. the dominant ions
279 occurring at m/z 360, 372, 374, and 390 (during desorption) corresponded to the iodide-cluster
280 ions $C_8H_{11}NO_7I^-$, $C_{10}H_{15}NO_6I^-$, $C_{10}H_{17}NO_6I^-$, and $C_{10}H_{17}NO_7I^-$ (Fig. 2a). These correspond to the
281 MCM species C727PAN and C731PAN, C923PAN, NLIMALOH and LIMALNO3,
282 NLIMALOOH, respectively.



283
284 **Figure 2.** Representative average mass spectrum for the desorption of SOA collected during the experiments:
285 Identification of ions detected in the (a) monomer region (m/z 340–440) and (b) dimer region (m/z 580–700). (c)
286 Relative intensities and positions of the two regions detected in all aerosol samples. Data was obtained from four 1-h
287 desorption cycles of $12\text{-}\mu\text{g m}^{-3}$ samples from a mixture with a N_2O_5 /limonene ratio of 2.4. The un-clustered (i.e., not
288 clustered with I⁻) m/z of each ion is $-127 m/z$.

289 Elevated signals of monomer ions (e.g., $C_{10}H_{15}NO_7$ (m/z 388), $C_{10}H_{15}NO_8$ (m/z 404),
290 $C_{10}H_{17}NO_8$ (m/z 406), and $C_{10}H_{15}NO_9$ (m/z 420)), which are absent from the list of expected

291 products of the mechanism, also occurred in this region. These non-MCM species contributed
292 significantly to the total organic monomer signal, and MCM species accounted for only $43.5 \pm$
293 3.2% of the total monomer signal of all experiments. One common feature of the monomers
294 without a match in MCM is that they contain a nitrogen atom and have an oxygen number higher
295 than 6, which is a range of compounds that is not represented explicitly in the MCM.

296 Monomers with progressively more oxygenated monomers of the general formula
297 $C_{10}H_{15}NO_x$ were detected for $x = 5-9$ i.e. $C_{10}H_{15}NO_5-C_{10}H_{15}NO_9$ with $C_{10}H_{15}NO_6$ being the
298 dominant species in both the aerosol and gas phase in most experiments. Ions with molecular
299 formulas containing two nitrogen atoms, for example, $C_{10}H_{16}N_2O_8$ (m/z 419) and $C_{10}H_{18}N_2O_8$
300 (m/z 421), were also detected (Fig. 2a). Limonene and its primary products reacted only with
301 NO_2 , NO_3^* , and HNO_3 , yielding molecules that are most likely di-nitrate species, with additional
302 functional groups.F

303 Similar to the highly oxygenated multi-functional species (HOMs) resulting from the
304 ozonolysis of monoterpenes (Ehn et al., 2014; Jokinen et al., 2015), including limonene, many of
305 the observed species could be classified as extremely low-volatility organic compounds (i.e.,
306 ELVOCs, which play a key role in SOA formation (Donahue et al., 2012). Observations
307 performed under ambient conditions during the 2013 Southern Oxidant and Aerosol Study
308 (SOAS) revealed the presence of highly functionalized particulate organic nitrates containing 6–8
309 oxygen atoms (Lee et al., 2016). In that work, these species constituted 3% and 8% of sub- μm
310 aerosol mass during daytime and nighttime hours, respectively, and exhibited a distinct diurnal
311 pattern, typically reaching peak concentrations between midnight and the early-morning hours.
312 The gaseous parent compounds were identified as monoterpenes, matching ions measured in their
313 laboratory study on α -pinene, enforcing the importance of monoterpene nitrates in the ambient
314 atmosphere. Complementary, Nah et al. (2016) also measured a large suite of highly oxygenated
315 organic nitrates from NO_3 oxidation of α -pinene and β -pinene in laboratory experiments.

316 For all elevated ion signals above m/z 390, there was no corresponding product in the
317 MCM mechanism. As shown in Fig. 2b, zooming into m/z 580-700 illustrating the high mass
318 dimer region, the largest ion signals corresponded to compounds with 19 and 20 carbons in the
319 dimer region. $C_{20}H_{22}N_2O_8$ and $C_{20}H_{29}NO_{17}$, which occurred at significantly elevated levels in all
320 aerosol samples, constituted the lowest- and highest-mass dimers, respectively (see Fig. 2 for
321 other examples of C_{19} and C_{20} dimer species). Many of these can be considered ELVOC species
322 based on their respective formulas and their partitioning behavior (i.e., they were present only in
323 the aerosol phase and at insignificant levels in the gas samples). $C_{19}H_{28}N_2O_x$ and $C_{20}H_{29}NO_x$ were
324 the most dominant families of C_{19} and C_{20} dimers, respectively. Taken together, 10 individual
325 dimers from these two families were identified in all experiments.

326 The contributions of the 11 most prevalent ion families (defined as groups of molecular
327 compositions with only the number of O atoms varying) to the total desorbed organic signal are

328 summarized in Table 2. Average contributions are calculated from the mean signals for each
329 family relative to the total mean organic signal generated during all experiments.
330
331

332 **Table 2.** Peak desorption temperature (T_{\max}) and the average contribution (over all experiments) to the organic signal
 333 during SOA desorption for the most commonly observed product families. The number of monomer species in each
 334 family that desorbed at only high temperatures is noted in parentheses.

Class	#	Family	# Observed in Family	Average Contribution	T_{\max} Range ($^{\circ}\text{C}$)
Monomers	m1	C10H15NO _x	5 (1)	23.0 ± 8.0 %	74 – 152
	m2	C10H18N2O _x	2 (0)	8.8 ± 2.4%	66 – 70
	m3	C10H16N2O _x	5 (1)	6.7 ± 2.2%	52 – 154
	m4	C10H17NO _x	5 (2)	5.3 ± 2.7%	59 – 159
	m5	C8H11NO _x	3 (0)	4.7 ± 1.4%	68 – 81
	m6	C9H13NO _x	4 (0)	3.0 ± 1.1%	70 – 75
	m7	C9H15NO _x	4 (0)	2.0 ± 0.7%	64 – 76
Dimers	d1	C20H29NO _x	4	7.1 ± 3.3%	100 – 154
	d2	C19H28N2O _x	6	5.0 ± 2.2%	101 – 157
	d3	C20H27NO _x	4	2.8 ± 1.2%	101 – 151
	d4	C20H24N2O _x	3	2.0 ± 1.7%	125 – 157

335

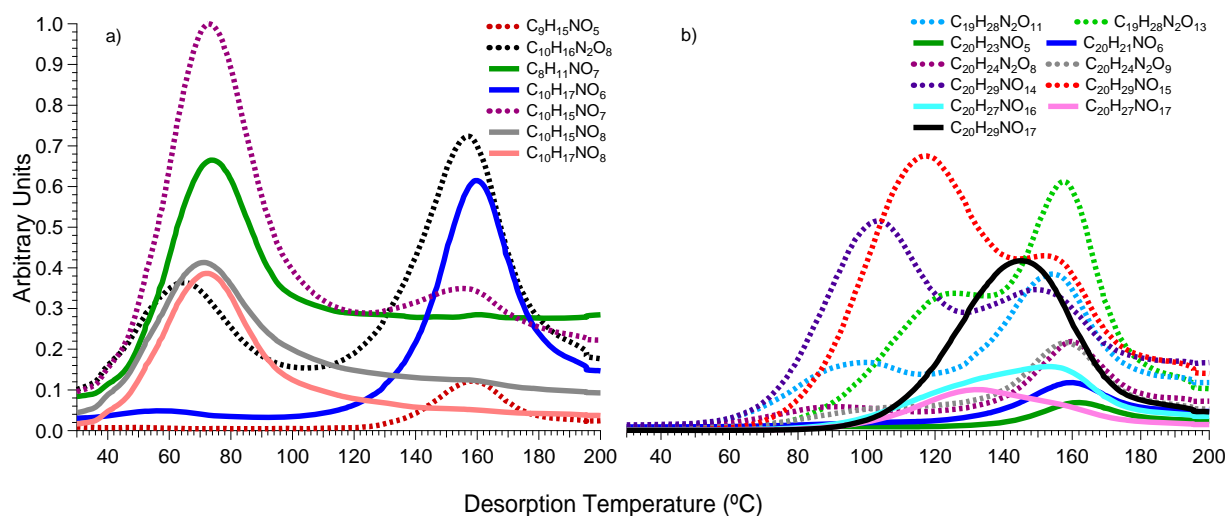
336 3.2 Characterization of identified ions via thermal properties

337 The desorption data is characterized by the frequent occurrence of multiple peaks
 338 corresponding to certain ions, and the thermograms in all experiments reveal four characteristic
 339 desorption patterns, which exhibit the following trends: (i) from 45 to 85°C, some monomer
 340 species undergo almost complete desorption. (ii) Some monomers yield two peaks - one in the
 341 low-temperature range and another at significantly higher temperatures. Additionally, (iii) some
 342 monomer ions, associated with certain individual species of the monomer families, occurred at
 343 only very high desorption temperatures, owing possibly to the fragmentation of high-mass
 344 oligomers and dimers. (iv) Although less prominent than that observed for monomers, a double
 345 peak occurred for several dimers, whereas for other dimers a single primary desorption peak
 346 occurred at mid to high temperatures (110–170°C). The occurrence of multiple peaks is
 347 consistent with the thermal degradation of extremely low-volatility species that desorb only at
 348 temperatures >200°C. Similar behavior has been observed in previous studies (Holzinger et al.,
 349 2010; Lopez-Hilfiker et al., 2014, 2015; Yatavelli et al., 2012), where the secondary peaks
 350 observed during desorption were attributed to the thermal degradation of very low-volatility
 351 aerosol components.

352 Analysis of the desorption profiles (thermograms) may yield additional information about
 353 the properties of each detected chemical species. The gradual heating of the FIGAERO filter
 354 from 25°C to 200°C resulted in a clear volatility-based separation of species and, for each ion
 355 detected, the desorption temperature corresponding to the maximum signal was identified.
 356 Furthermore, the average desorption temperature of the monomer species was typically lower
 357 than that of their dimer counterparts, which are less volatile. Higher masses (than those

358 associated with the monomer species) were typically desorbed from the FIGAERO filter at higher
 359 temperatures. An example of this characteristic behavior is shown in the average thermograms
 360 (Fig. 3) of several monomer and dimer ions. In general, compounds evaporating at relatively low
 361 temperatures were also found in the gas phase, indicative of monomer that partitioning between
 362 gas and particle phase.

363
 364



365
 366

367 **Figure 3.** Average thermograms (over four desorption cycles) for an N_2O_5 ratio of 2.4. Thermograms of ion clusters
 368 of the (a) monomer species (C_8 – C_{10}) and (b) dimer (C_{19} – C_{20}) species. Ions with double-peak thermogram shape
 369 patterns, consistent with the fragmentation of low-volatility oligomers, are shown as dashed lines.

370

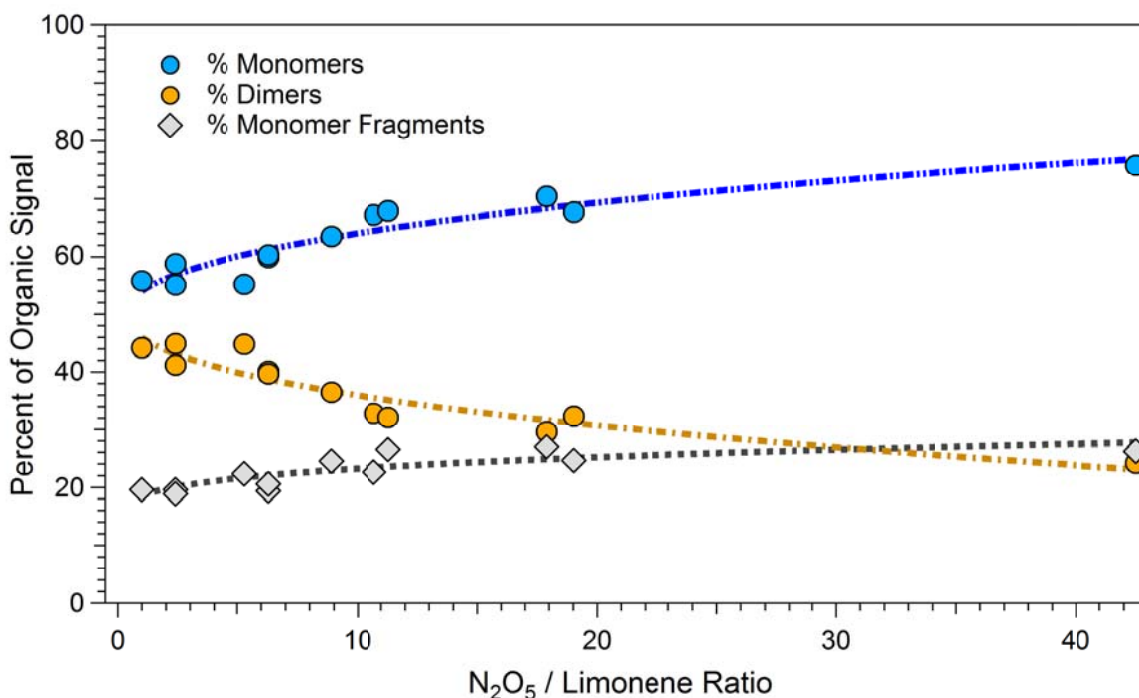
371 As shown in Fig. 3, each of the detected ion signals reaches at least one local maximum
 372 value. The temperature at which a signal reached the first maximum (T_{max}) value was similar
 373 across all experiments (average standard deviation: $<10\%$). Secondary peaks occurred more
 374 frequently for species with a carbon number of 10 or lower, consistent with a degradation-based
 375 contribution. Although the temperature at which the secondary local maximum occurs ($T_{\text{max},2}$)
 376 provides insight into the occurrence of dimerization, the T_{max} value was taken as the true
 377 desorption temperature of each ion.

378 T_{max} values were identified for each ion in the 196-ion set. Monomer, i.e., lower-mass,
 379 species ($\text{C} \leq 10$) desorbing at high temperatures could be produced as fragments via thermal
 380 degradation of higher-MW species. Some of these ions are matching the chemical composition
 381 ($\text{C}_{10}\text{H}_{16}\text{O}_4$, $\text{C}_{10}\text{H}_{17}\text{NO}_5$, $\text{C}_{10}\text{H}_{17}\text{NO}_6$, and $\text{C}_7\text{H}_{10}\text{O}_4$) of primary products within the MCM,
 382 accounting for (on average) $69.0 \pm 10.8\%$ of the signal detected in the gas phase. Here some

383 possibilities are plausible, one could be that they are produced as monomer but are important
384 building blocks in the dimer formation, thus thermally decompose back to monomers during
385 desorption.

386 The ratio of dimer to monomers varied between experiments. At high ratios of N_2O_5 to
387 limonene, the fraction of dimer species decreased relative to the total organic signal, whereas the
388 percentage of high-temperature desorbing monomer species (fragments) increased (Fig. 4). This
389 suggests that absolute dimer formation may have remained the same, but the monomer signal is
390 over-represented by monomer fragments generated from high-mass, thermally unstable
391 compounds. This percentage is calculated based on the assumption of a common detection
392 sensitivity across all ions; this assumption may influence the estimated (percentage) contribution
393 of monomers relative to that of dimers.

394



395
396 **Figure 4.** Percentage of monomer, dimer, and high-temperature monomer signal (observed during desorption)
397 relative to the ratio of N_2O_5 to limonene injected into the reactor. The data points at a ratio of 113 are not shown (22,
398 78, 39%, respectively). The lines indicated are for the guidance of the eye.

399

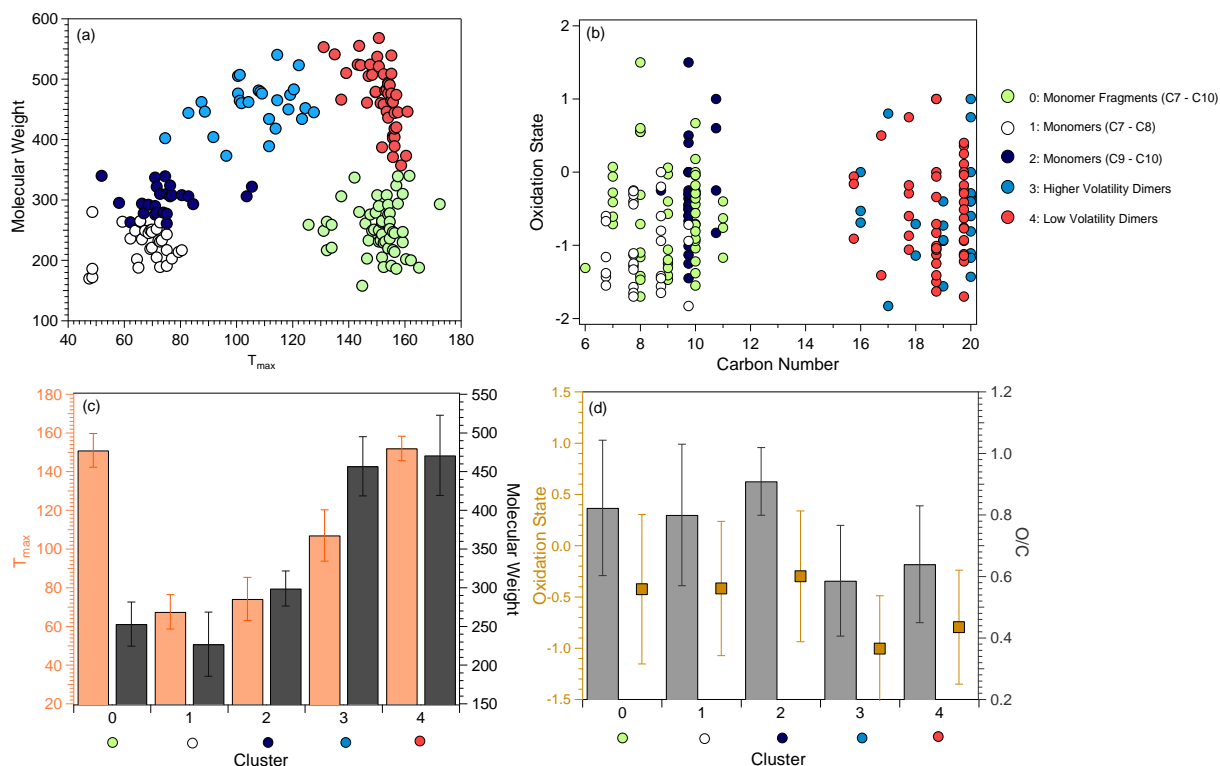
400 3.3 Characterization of major SOA products via cluster analysis

401 Clustering was performed on an ion set consisting of 117 ions, which accounted for >90%
402 of the total organic signal generated during desorption in all experiments. Ions generating
403 extremely low signal (i.e. the thermogram did not exhibit any structure identifiable above

404 background noise prohibiting Tofware to constrain a mathematical fit for T_{\max} calculations) were
405 excluded to prevent analysis of ions with mis-identified T_{\max} values. However, the occurrence of
406 high-temperature desorbing monomer outliers (described previously) and the double-peak
407 behavior exhibited by several monomers rendered the mass- and temperature-based grouping of
408 these ions difficult. To address this issue, duplicate entries, corresponding to T_{\max} and $T_{\max,2}$,
409 were assigned to all ions exhibiting double-peak behavior, allowing the clear separation and
410 analysis of low-mass ions desorbing at temperatures $>120^{\circ}\text{C}$.

411 Four and five clusters ($\#_{\text{clust}} = \{4, 5\}$), using T_{\max} , MW, and #C as input, yielded the best
412 T_{\max} -based clustering and separation of ions. The use of n_{H} and n_{O} as additional input parameters
413 resulted in partial separation of clusters into groups with similar O/C and H/C ratios, and poor
414 correlations with respect to T_{\max} . The average silhouette score obtained for four clusters was
415 better (0.81 vs. 0.72) than that obtained for five clusters. However, the use of five clusters
416 allowed for the separation of low-temperature desorbing monomers into two groups with distinct
417 average T_{\max} , and MW with smaller differences in O/C ratios, and oxidation states ($2\times\text{O/C} - \text{H/C}$
418 $- 5\times\text{N/C}$). Using more than five clusters resulted in a further decrease in the quality of *cluster*
419 *separation*, as measured by the inertia (Eq. 1) and average silhouette score (Eq. 2). Although the
420 identification of subgroups within each cluster are possible by increasing $\#_{\text{clust}}$, the five main
421 clusters were chosen based on their separation by mass and T_{\max} values and to reduce complexity
422 of the interpretation of the resulting clusters with respect to the chemical composition.

423 Figure 5a shows the cluster separation on the MW– T_{\max} plane. The distribution of
424 individual cluster members based on oxidation states and #C (Fig. 5b), and the mean MW, T_{\max} ,
425 O/C, and oxidation state of each cluster (Fig. 5c) are also shown.



426
 427 **Figure 5.** Characteristics of the five identified clusters: (a) Desorption temperature of each observed ion in the top
 428 40th percentile of ions (identified by their respective desorption signal), color-coded by their corresponding cluster
 429 number, (b) Oxidation state relative to carbon number of all observed ions, colored by their corresponding cluster
 430 membership (for visualization purposes, carbon numbers of groups 0, 2, and 4 are offset), (c) average cluster mass
 431 and desorption temperature, and (d) average cluster oxidation state and O/C ratios. Error bars in panels (c) and (d)
 432 indicate standard deviations for each cluster property.

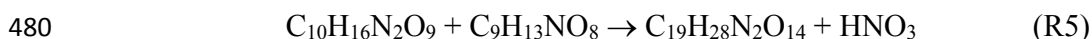
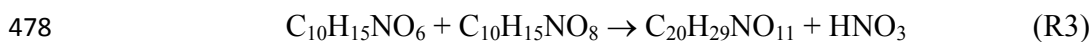
433
 434 As Fig. 5 shows, the five clusters are characterized by distinct average MWs and
 435 corresponding average T_{max} values. Cluster 0 consists of monomer ions, which are considered
 436 fragments of larger, less-volatile molecules that desorb at high temperatures. The average
 437 oxidation state and O/C ratio are similar to those of clusters 1 and 2, which are composed
 438 primarily of C7–C9 and C9–C10 monomer ions, respectively. This results from the fact that 87%
 439 and 69% of cluster 1 and 2 ions, respectively, have secondary thermogram peaks and T_{max} values,
 440 and the ions represented as members of both clusters 1 and 0. Ions corresponding to the identified
 441 dimers are contained in clusters 3 and 4. The dimers are characterized by two primary desorption
 442 regimes, with species that desorb at mid-range temperatures (80–130°C) occurring in cluster 3
 443 and the highest-mass, lowest-volatility ions occurring in cluster 4. Moreover, the distribution of
 444 individual cluster members with respect to #C and oxidation state (Fig. 5b) shows that members
 445 of low-MW clusters (0, 1, 2) and high-MW clusters (3, 4) reside in separate regimes. The ions in
 446 high-MW clusters have a significantly larger number of carbon atoms per molecule and, hence,
 447 lower (on average) oxidation states than ions in clusters 0–2. With respect to the most prevalent

448 families listed in Table 2, monomer families m2, m3, and m4 reside exclusively in cluster 2,
449 whereas m5 and m7 reside exclusively in cluster 1. Family members of m1 and m6 were split
450 20/80% and 75/25% between clusters 1 and 2, respectively. Dimer families d1–d4 occurred
451 predominantly (66–75%) in cluster 4, with the remainder residing in cluster 3. None of the dimer
452 families in Table 2 occurred in clusters 0, 1 or 2.

453 A positive trend between the Mw and Tmax values, see Fig 5a., was obtained for data in
454 two of the monomer clusters (1 and 2) and the high volatile dimer cluster, while the trend turned
455 negative for the low volatile dimers cluster. It should be noted that monomer species had (in
456 general) higher O/C ratios than the dimers. It could be that monomers need more oxidation before
457 being transferred into the condensed phase. However, as outlined by the partitioning plots (Fig.
458 S1) most monomers also have a significant condensed phase contribution. Rather, this
459 observation provides some insight into the processes of dimerization that are occurring,
460 indicating the extent to which oxygen is lost during the dimerization process.

461 3.4 Mechanisms of dimerization

462 The mechanism to create dimers with one nitrogen and a lower O/C ratio would
463 presumably involve the loss of a nitrogen oxides or nitric acid. For this complex system and
464 within the scope of this study it was not possible to firmly proof any mechanism. Since the
465 experiment were done at low RH the direct hydrolysis would be less likely (see Rindelaub et al,
466 2015, 2016). However, knowing HNO₃ being thermodynamic stable one may speculate in that
467 dimerization of two monomer species via the loss of one HNO₃ molecule could occur e.g. where
468 a C₂₀H₂₉NO_y (y = 7–15) species would be generated from C₁₀H₁₅NO_x (x = 5–9) species. This
469 process could be seen as the reverse of esterification in order to produce a dimer product with one
470 less nitrogen and reduced numbers of oxygens. For example, with HNO₃ as a leaving group, the
471 mechanism of dimerization between C₁₀H₁₅NO₆ and C₁₀H₁₅NO₈ (see Reaction 3), would produce
472 the C₂₀ dimer species (C₂₀H₂₉NO₁₁) that was observed in all experiments. The formation of the
473 observed C₁₉ dimer species (e.g., C₁₉H₂₇O₁₅) through the combination of, for example,
474 C₁₀H₁₇NO₇ and C₉H₁₁NO₁₁ monomer species (Reaction 4) is also attributed to this mechanism.
475 Additionally, the occurrence of dimer species with two nitrogen atoms, through the combination
476 of monomers such as C₁₀H₁₆N₂O₉ and C₉H₁₃NO₈ (Reaction 5), can also be attributed to this
477 dimerization mechanism.



481 The higher O/C ratios of the monomer species, compared with those of the
482 dimers/oligomers, may also be attributed to the loss of an HNO₃ molecule (from the monomer)
483 during the dimerization process. For example, the two C₁₀ reactants in Reaction 3 have O/C ratios
484 of 0.6 and 0.8 while the product, C₂₀H₂₉NO₁₁, has an O/C ratio of 0.55. A similar trend is
485 observed for Reactions 4 and 5, where the reactants have an average O/C ratio of 0.96 and 0.89,

486 respectively, and the products have O/C ratios of 0.79 and 0.74, respectively. Due to the loss of
487 HNO₃ during dimerization, the potential dimer decomposition during desorption is expected to
488 yield fragments which differ in molecular composition from the precursor (i.e., pre-dimerization)
489 monomers. However, the resulting monomers may also be associated with aerosol phase products
490 that have secondary desorption peaks. For example, the fragmentation of C₂₀H₂₉NO₁₁ could yield
491 C₁₀H₁₄O₆ + C₁₀H₁₅NO₅ or C₁₀H₁₆O₅ + C₁₀H₁₃NO₆, and the fragmentation of C₁₉H₂₇NO₁₅ might
492 yield C₉H₁₃NO₆ + C₁₀H₁₄O₉ or C₉H₁₃NO₉ + C₁₀H₁₄O₆. Likewise, C₉H₁₃NO₇ + C₁₀H₁₅NO₇ or
493 C₉H₁₃NO₉ + C₁₀H₁₅NO₅ monomer pairs could be generated from the thermal degradation of
494 C₁₉H₂₈N₂O₁₄.

495 The fragmentation of dimers may also proceed through multiple channels, thereby
496 producing several sets of monomer fragments, or the fragmentation of multiple dimers may
497 produce the same ions. Therefore, attributing the production of a monomer fragment to the
498 thermal degradation of a specific dimer is difficult, using the current dataset. Large (C > 20)
499 oligomeric species may contribute to the high-temperature generation of monomer fragment
500 species. The proposed mechanisms may play only a partial role in the dimerization process
501 occurring in these experiments. However, they offer a plausible explanation for the occurrence of
502 multiple observed dimers and the secondary desorption maxima associated with the monomer
503 constituents.

504 **4 Conclusions**

505 High-resolution mass spectrometric data was analysed for condensed-phase reaction
506 products resulting from NO₃ initiated oxidation of the monoterpene, limonene. The results
507 revealed that the formation of organic nitrates contributed substantially (89.5 ± 1.4% of the
508 particulate-phase ion signal) to SOA formation, with dimers constituting a significant fraction of
509 the particle-phase products. On average, monomers and dimers/oligomers contributed 63 ± 7 and
510 37 ± 7%, respectively, of the particle-phase organic signal detected by the I-CIMS. Furthermore,
511 many monomers (accounting for 22 ± 3% of the average organic signal) desorbed at high
512 temperatures (120°C). The fraction of the signal generated by monomers increased with
513 increasing N₂O₅/limonene ratio (ratio of 43 yields a fraction of 76%), whereas the fraction of
514 dimers decreased (to 24%). The fraction of the monomer signal resulting from desorption at high
515 temperatures (≥120°C) also increased (by 26%). Therefore, although the monomer fraction
516 increased with increasing N₂O₅/limonene ratio, this increase in desorption signal occurred
517 primarily at temperatures above 120°C, indicative of an increase in the fragmentation of high-
518 MW dimers and oligomers. A large portion (79%) of the monomer thermograms exhibited this
519 bi-modal behavior, with secondary peaks occurring above 120°C, indicating that the composition
520 of SOA was largely determined by the formation of thermally unstable, low-volatility oligomers.

521 In total, 196 individual organic ions were detected during desorption. However, the total
522 measured organic signal was generated mainly by 52 (i.e., 76%) of these ions, which constituted
523 the 75th percentile of the monomer and dimer signals. Over half of the signal emanated from the

524 top 90th percentile, which comprised a small subset of only 20 species, of the total number of
525 ions. These 20 species (with nine listed as major products in the MCM) constituted the major
526 particle-phase products formed via the reaction of N₂O₅ and limonene under the conditions
527 employed in this study. The non-listed species (see Table S1) were either dimer species or more
528 highly oxygenated, nitrated analogs of known major products, which are notoriously hard to
529 describe via standard gas-phase mechanisms. There are two frequently suggested pathways for
530 these. Firstly, the high number of oxygens would be result of isomerization of RO or RO₂ that
531 rarely is described explicit in current modelling framework. Secondly, the presence of di-nitrated
532 compounds relies on secondary chemistry derived from e.g. produced mononitrates
533 intermediates; for limonene containing two double bonds this is more relevant than for other
534 monoterpenes and so far not commonly described in models.

535 Cluster analysis revealed two monomer groups, two dimer groups and, a separate group
536 containing monomer ions that exhibited secondary desorption peaks occurring at temperatures
537 $\geq 150^{\circ}\text{C}$. Each group was characterized by a distinct average MW and desorption temperature
538 (T_{max}). The 2 identified clusters in the monomer and dimer sub-classes differ in oxidation state
539 and O/C ratios, with increasing O/C corresponding to higher T_{max} values.

540 Using a combination of cluster analysis and thermal properties derived from FIGAERO-
541 CIMS measurements may provide some means of reducing the complexity associated with the
542 description of SOA formation processes. The investigated reaction system constitutes only one of
543 many systems, but could be used as an example of the evaluation required for this type of
544 information derived from high-resolution MS. The results revealed that, analogous to products
545 from ozonolysis and $\cdot\text{OH}$ -induced oxidation, the organic nitrates produced in the nighttime
546 chemistry of biogenic compounds comprise a multi-component mixture that contributes to
547 ambient SOA. Thus, the aerosol species detected here could be included in modeling studies with
548 the aim of explaining scenarios where SOA formation rates are under-predicted. Furthermore, the
549 numerous products resulting from NO₃ oxidation of limonene, which were identified and grouped
550 based on thermal properties, could be candidates for identification in ambient air masses
551 dominated by nocturnal limonene chemistry.

552 **Competing interests**

553 The authors declare that they have no conflict of interest.

554

555 **Acknowledgment**

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