

Point-by-point response to review comments on manuscript acp-2017-582 “Quantifying black carbon light absorption enhancement by a novel statistical approach”

By Cheng Wu et al.

We thank the two anonymous reviewers for their comments to improve the manuscript. Our point-by-point responses to the review comments are listed below. Changes to the manuscript are marked in blue in the revised manuscript. The marked manuscript is submitted together with this response document.

Anonymous Referee #2

R2-Q1. The authors have made a number of improvements to the manuscript and adequately addressed my original comments with two important exceptions, detailed below. If these can be addressed to the satisfaction of the editor I recommend publication in ACP.

The response to comments regarding the influence of RH on the Aethalometer measurements is not adequate. The response states that a previously observed strong correlation between a non-dried Aethalometer and a dried photoacoustic instrument remained high even under elevated RH conditions. First, a strong correlation does not mean one of the instruments is not biased. For example, an instrument biased by a constant factor of two will still be strongly correlated with a reference measurement. A bigger point is that the data shown in the response show no increase in absorption measured by the Aethalometer relative to a dried reference measurement for RH values approaching 100%. This means that $f(\text{RH})$ is flat (1.0) for the Guangzhou data and there is no LWC influence on absorption. There could be real reasons for a difference in $f(\text{RH})$ during the two studies, but if this is the case, the Guangzhou data cannot be used to validate the response of the Aethalometer, since the conditions are different. Until the response of the Aethalometer has been validated against a reference instrument at high relative humidity conditions and at a location where there is an RH impact on absorption, Aethalometer data alone has too many uncertainties to draw the types of conclusions made in Section 4.5.

Author’s Response: We understand the concerns regarding to the RH impact on Aethalometer measurements. RH effect is challenging to account for with σ_{abs} measurement by in-situ techniques including PAS and extinction-minus-scattering approach. Since the $f(\text{RH})$ of aerosol light absorption is a much smaller quantity compared to that of scattering, the instrument precision is crucial for the determination of light absorption $f(\text{RH})$. Considering that we do not have data to validate the response of Aethalometer against a reference instrument at high RH conditions, we now have removed section 4.5 in the revised manuscript. We are looking forward to the future advancement of σ_{abs} instrumentation to provide more precise measurements under high RH. $f(\text{RH})$ of light absorption is an important factor that needs to be investigated in the future, when the proper instrumentation become available.

The point we would like to make by showing the PAS-Aethalometer comparison dataset is that, the non-dried Aethalometer measurement correlated well with a dried PAS measurement. We do not intend to imply that one of the instrument is not biased.

While we have decided note to include in this manuscript the discussion on the dependency of optical properties on RH, we would like to address the below question raised by the reviewer: if $f(\text{RH})$ does exist in the Aethalometer measurement, will it affect the correlation between a non-dried Aethalometer and a dried PAS? We conducted a test using synthesized data to answer the question and the details are described below.

First, σ_{abs} by a dried PAS is generated (Figure R1a) following a lognormal distribution using Mersenne Twister (MT) pseudorandom number generator. Then σ_{abs} by a dried Aethalometer is generated (Figure R1b), which is highly correlated with σ_{abs} of a dried PAS (Slope=1, $R^2=0.99$).

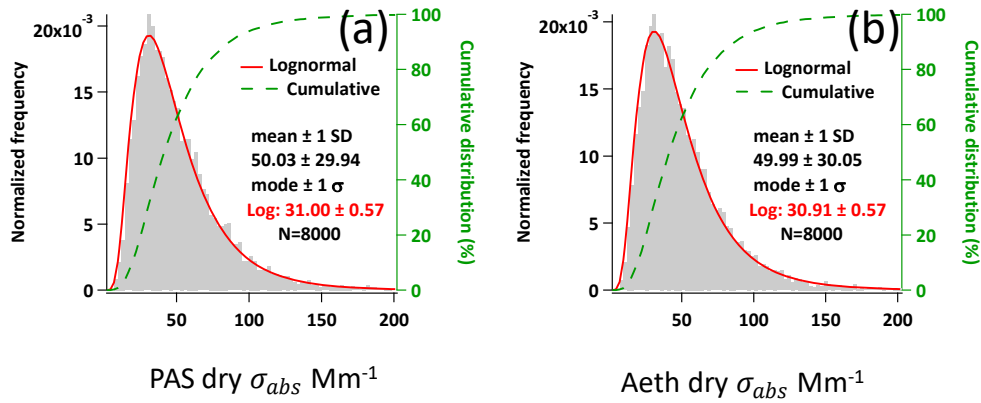


Figure R1. Distribution of synthesized σ_{abs} . (a) Dried PAS and (b) Dried Aethalometer

Then RH is generated independently following a uniform distribution in the range of 30~100% as shown in Figure R2a. The corresponding non-dried Aethalometer σ_{abs} is derived as shown in Figure R2c following the $f(RH)$ (distributions are shown in Figure R2b).

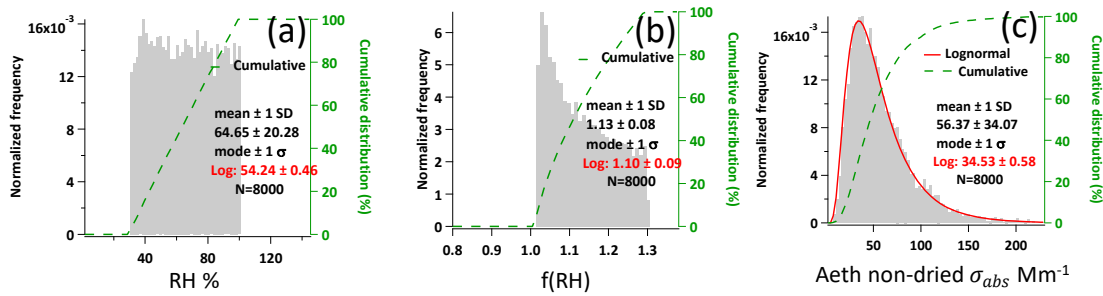


Figure R2. Distribution of synthesized data. (a) RH; (b) $f(RH)$; and (c) non-dried Aethalometer σ_{abs} .

Scatter plots with linear regression lines are shown in Figure R3. From the simulation perspective, the inclusion of $f(RH)$ only decrease the R^2 by a small amount (from 0.99 to 0.97). As a result, the high R^2 (0.96) alone shown in Figure S1 does not necessarily lead to the conclusion that $f(RH)$ is flat, instead, the slope indicates whether $f(RH)$ is >1 or flat. Since the Dried Aethalometer vs. Dried PAS is not tested with collocated non-dried Aethalometer vs. Dried PAS in our previous field measurements, the $f(RH)$ effect requires further investigation in the future.

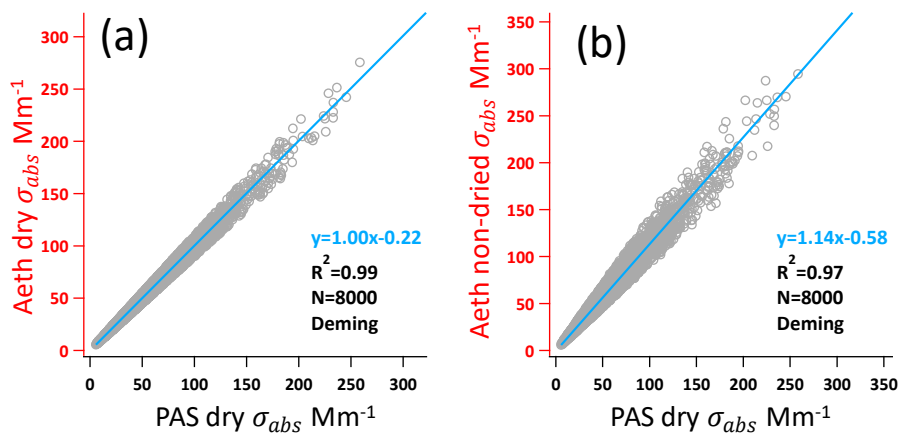


Figure R3. Scatter plots of synthesized data. (a) Dried Aethalometer vs. Dried PAS. (b) non-dried Aethalometer vs. Dried PAS.

R2-Q2. The addition of the tests to investigate impact of bias on the results is helpful, however my original comments were meant to question effects of time-varying biases on the data that depended on the amount of “coating” present or other physical processes not related to the real enhancement of light absorption. I apologize for not being more clear. The authors should generate a test data set with varying EC and prescribed “coating” contributions to measured absorption and EC, then introduce a bias to both data sets that depends on the amount and/or contribution of coating. Also useful would be tests or demonstrations of how other phenomena could affect the MSR method (e.g., a period of dust impact that affects absorption but not EC, or affects EC through mineral catalysis of EC, evaporation and/or adsorption of semi-volatile material from and/or onto filters in both the Aethalometer or OC/EC method). I believe this is merited given the novelty of the approach and analysis.

Author’s Response: To further address the impact of measurement biases, we have re-organized section 5 (Caveats of the MRS method in its applications to ambient data) into four parts, covering the following aspects:

- 5.1 Impact of measurement biases
- 5.2 Impact of semi-volatile organic carbon
- 5.3 Impact of mineral dust
- 5.4 Impact of BrC

The discussions on systematic bias (Test A&B) are moved from section 4.1 to section 5.1 (Impact of measurement biases). Test C is introduced in section 5.1 to investigate the impact of sample-dependent bias as a function of E_{abs} as shown below.

Study by Cheng et al. (2016) found two distinct types of biomass smoke behave differently on the biases of filter based σ_{abs} measurement. The bias in the first type can be explained by a nearly constant correction factor, which is similar to the situation discussed in Test A. The bias in the second type shows an apparent OC/EC dependence. Test C is carried out to investigate this situation, i.e., examining the impact of sample-dependent bias as a function of E_{abs} . Unlike the proportional bias in Test A and B that is the same for all data points, the bias in Test C depends on the E_{abs550} of individual samples, which are parametrized by Eqs. (11) and (12).

$$\sigma'_{abs550} = \sigma_{abs550} + \sigma_{abs550} \times (k \times E_{abs550} - k) \quad (11)$$

$$EC' = EC - EC \times (k \times E_{abs550} - k) \quad (12)$$

As shown in Eqs. (11) and (12), the positive bias of σ_{abs550} and negative bias of EC are proportional to E_{abs550} . The magnitude of E_{abs550} -dependent bias is regulated by the factor k . Since σ'_{abs550} and EC' are biased in different directions, resulting a further amplification in MAE biases, which could be considered as the extreme case. As shown in Figure S20, for $k=10\%$ (corresponding to a bias of 10% when $E_{abs}=2$), the bias of MRS-derived E_{abs} is very small (1%). For $k=20\%$, the MRS-derived E_{abs} changes from 1.44 to 1.66, leading to a bias of 15%. These results imply that if the measurement bias follows the same form as demonstrated in Test C, the bias is not negligible but still acceptable. If the impact only affects σ_{abs} or EC rather than impacting both, the bias is expected to be smaller than the estimation shown in Test C.

It should be noted that the parameterization scheme shown in Eqs. (11) and (12) is only for demonstration purpose from a conceptual perspective and it does not necessarily represent the real-world measurements. There is a lack of quantitative understanding of this impact. For example, Lee et al. (2007) used artificially fabricated EC samples with OC coatings to evaluate the impact of coating on OC/EC analysis.

Biases were observed, nevertheless, the results were linearly correlated with the true OC and EC values with a high R^2 (>0.9), implying that the biases in that specific study were dominated by systematic biases rather than the coating-dependent bias. Further studies are needed to better characterize and parameterize this impact if filter-based techniques are used for σ_{abs} and EC determination in the MRS approach.

To address the impact of semi-volatile material, following contents are added as section 5.2 (Impact of semi-volatile organic carbon)

5.2 Impact of semi-volatile organic carbon

Light absorption contribution due to semi-volatile organic carbon (SVOC) from wood combustion was reported to be negligible in the visible range and around 10-20% at 360 nm (Chen and Bond, 2010). On the other hand, OCEC analysis can be affected by SVOC (Subramanian et al., 2004), either by positive artifacts through adsorption of SVOC onto quartz filters, or by negative artifacts through evaporation of SVOC due to the gas-particle re-equilibrium down stream of VOC denuder. Positive artifacts can be minimized by the installation of a VOC denuder which is widely adopted in RT-OCEC measurements (Bae et al., 2004; Bauer et al., 2009). A typical negative artifact of 10% is expected and can be corrected by backup filters (Subramanian et al., 2004). There was evidence to show that SVOC could affect OC/EC split in thermal/optical analysis (Cheng et al., 2009). However, the bias in EC caused by the OC/EC split drift due to SVOC is systematic, making it falls into the scenario discussed in Test B. As a result, the impact from SVOC on E_{abs} estimation by MRS is expected to be small.

Section 5.3 is added to provide the recommendations for samples strongly influenced by mineral dust.

5.3 Impact of mineral dust

The presence of mineral dust (MD) could affect both σ_{abs} and EC determination. If MD is externally mixed with soot particles, the light absorption from MD could be miscounted as σ_{abs} enhancement, leading to the overestimation of E_{abs} . If the light absorption signal from MD is sufficiently strong (e.g. $AAE > 2$), σ_{abs} by MD and BC can be separated by the AAE approach suggested by Fialho et al. (2005). Additionally, the presence of substantial MD in samples has several impacts on the EC determination by thermal analysis. First, if the samples are not pre-treated with acid, the carbonated carbon could be misidentified as EC, resulting over-estimation of EC (Chow et al., 1993). The acid treatment is only available for off-line OC/EC analysis and not yet practical for the RT-OCEC analyzer. Second, metal oxides in MD can lead to premature EC oxidation in the helium stage of OC/EC analysis, leading to underestimation of EC (Wang et al., 2010; Bladt et al., 2012). The lack of a parameterization scheme for correcting the EC loss due to MD makes it improper to

use the biased EC as a primary tracer. For these reasons, E_{abs} estimation by MRS is not recommended for samples strongly influenced by MD.

Discussion on the impact of BrC (added in the last revision) is now moved to section 5.4.

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1 **Quantifying black carbon light absorption enhancement by a novel**
2 **statistical approach**

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18 **Abstract**

19 Black carbon (BC) particles in the atmosphere can absorb more light when coated by non-
20 absorbing or weakly absorbing materials during atmospheric aging, due to the lensing effect. In this
21 study, the light absorption enhancement factor, E_{abs} , was quantified using one year's measurement of
22 mass absorption efficiency (MAE) in the Pearl River Delta region (PRD). A new approach for
23 calculating primary MAE (MAE_p), the key for E_{abs} estimation, is demonstrated using the Minimum R
24 Squared (MRS) method, exploring the inherent source independency between BC and its coating
25 materials. A unique feature of E_{abs} estimation by the MRS approach is its insensitivity to systematic
26 biases in EC and σ_{abs} measurements. The annual average $E_{\text{abs}550}$ is found to be 1.50 ± 0.48 (± 1 S.D.)
27 in PRD, exhibiting a clear seasonal pattern with higher values in summer and lower in the winter.
28 Elevated E_{abs} in the rainy summer season is likely associated with aged air masses dominating from
29 marine origin, along with long-range transport of biomass burning influenced air masses from
30 Southeast Asia. Core-shell Mie simulations along with measured E_{abs} and Angstrom absorption
31 exponent (AAE) constraints suggest that in the PRD, the coating materials are unlikely to be dominated
32 by brown carbon and the coating thickness is higher in the rainy season than the dry season.

33 **1 Introduction**

34 Originating from incomplete combustion, black carbon (BC) is a crucial constituent of
35 atmospheric aerosols, and is an air pollutant itself, having an adverse health impacts on humans (Suglia
36 et al., 2008). BC has also been recognized as the third most important climate forcer due to its broad
37 light absorbing capability across the UV-Vis-IR spectrum (IPCC, 2013). BC can alter the climate in a
38 variety of ways, including by direct forcing (Bond et al., 2011), affecting cloud cover (Koch and Del
39 Genio, 2010) and precipitation (Tao et al., 2012), reducing the albedo of snow and ice (Hansen and
40 Nazarenko, 2004) and causing surface dimming (Wild, 2011). The climate effects of BC can be global
41 or regional (Ramanathan and Carmichael, 2008). A recent study found BC can modify planetary
42 boundary layer meteorology, and thus enhance local pollution indirectly (Ding et al., 2016). However,
43 due to its variable optical characteristics induced during atmospheric aging, large uncertainties still
44 exist in estimating the radiative forcing from BC. Optical properties of BC can be predicted by
45 knowing the mass concentration, mixing state and size distribution, which collectively serve as the
46 cornerstone for modeling the climate effect of BC. In 3D modeling studies, to conserve computational
47 resources, the mass absorption efficiency (MAE) or mass absorption cross-section (MAC) is widely
48 used to convert black carbon mass concentration to light absorption coefficient (σ_{abs}). MAE is a
49 quantity to describe the light absorption ability per unit EC mass:

$$50 \quad \text{MAE } (m^2 g^{-1}) = \frac{\text{absorption coefficient } \sigma_{abs} (Mm^{-1})}{\text{EC mass concentration } (\mu g m^{-3})} \quad (1)$$

51 As a fundamental input parameter, MAE has a critical impact on BC's radiative forcing
52 estimation in climate modeling studies. Mixing state is one of the governing factors affecting MAE.
53 Light absorption of soot particles is enhanced when coated with non-absorbing materials (Fuller et al.,
54 1999) or weakly absorbing materials (Lack and Cappa, 2010) during atmospheric aging. The coating
55 materials can focus more light onto the soot core through the lensing effect, resulting in elevated MAE
56 (Wang et al., 2017). Strong correlations between MAE and the number/volume fraction of coated
57 particles have been reported in urban areas like Tokyo (Naoe et al., 2009), Shenzhen (Lan et al., 2013)
58 and Xi'an (Wang et al., 2014), implying that the elevated MAE observed at these locations was mainly
59 due to the elevated fraction of coated of soot particles. Total absorption ($\sigma_{abs,t}$) of coated particles can

60 be separated into two parts: primary absorption ($\sigma_{abs,p}$) due to the uncoated soot core alone, and extra
 61 absorption ($\sigma_{abs,c}$) due to lensing effect of the coating (Bond et al., 2006; Jacobson, 2006; Liu et al.,
 62 2016a) and the presence of secondarily formed brown carbon (BrC) (Lack and Cappa, 2010; Liu et al.,
 63 2016b).

$$64 \quad \sigma_{abs,t} = \sigma_{abs,p} + \sigma_{abs,c} \quad (2)$$

65 The absorption enhancement factor (E_{abs}) then can be defined as ratio of the total absorption and
 66 primary absorption coefficients or the corresponding MAE values:

$$67 \quad E_{abs} = \frac{\sigma_{abs,t}}{\sigma_{abs,p}} = \frac{MAE_t}{MAE_p} \quad (3)$$

68 Where MAE_p represents the ratio of $\sigma_{abs,p}/EC$ for uncoated soot particles, similar to the concept of
 69 the primary OC/EC ratio in the EC tracer method:

$$70 \quad MAE_p = \frac{\sigma_{abs,p}}{EC} \quad (4)$$

71 And the MAE of coated BC can be defined as:

$$72 \quad MAE_t = \frac{\sigma_{abs,t}}{EC} \quad (5)$$

73 Thus, elevated MAE induced by coating during atmospheric aging results in an E_{abs} larger than 1.

74 Previous model studies suggest that absorption by aged soot particles can be 1.5 times greater than
 75 fresh soot (Fuller et al., 1999; Bond et al., 2006). Laboratory studies have demonstrated that soot
 76 particles coated with SOA (Saathoff et al., 2003; Schnaiter et al., 2005; Tasoglou et al., 2017) and
 77 sulfuric acid (Zhang et al., 2008; Khalizov et al., 2009) can increase E_{abs} . An artificial coating
 78 experiment by Shiraiwa et al. (2010) found an E_{abs} of 2 for graphite particles growing in diameter from
 79 185 to 370 nm. A laboratory study by McMeeking et al. (2014) found that in the presence of BrC, light
 80 absorption enhancement is more pronounced at the shorter wavelength. A recent chamber study
 81 coupling actual ambient air with seed BC particles implies that the timescale for E_{abs} reaching 2.4 is
 82 only 5 hours in Beijing but 18 hours in Houston (Peng et al., 2016). Field studies conducted in recent
 83 years have also substantiated enhanced light absorption in Canada (Knox et al., 2009; Chan et al.,
 84 2011), US (Lack et al., 2012b), UK (Liu et al., 2015) and Japan (Nakayama et al., 2014; Ueda et al.,

85 2016). In contrast, field studies in California, US (Cappa et al., 2012) found a weaker light absorption
86 enhancement (6% on average). A recent study suggests the mass ratio of non-BC content to BC
87 particles determines the occurrence of the absorption enhancement of black-carbon particles (Liu et
88 al., 2017).

89 Two approaches are widely used to determine E_{abs} from ambient measurements. The first approach
90 removes the coating materials on particles physically using a thermal denuder (TD) (Lack et al., 2012a)
91 or by aerosol filter filtration-dissolution (AFD) (Cui et al., 2016b). The TD approach is briefly
92 discussed here. Coating materials can be removed by TD at a working temperature around 200 to
93 300 °C (depending on the charring characteristics of aerosols at the sampling site) to measure $\sigma_{abs,p}$,
94 which are cycled with measurements of $\sigma_{abs,t}$ (without passing through TD), allowing E_{abs} to be
95 obtained from the ratio of $\sigma_{abs,t}/\sigma_{abs,p}$ following Eq.3. The major advantage of the TD approach is
96 its ability to provide highly time resolved measurements (minutes). A photo-acoustic spectrometer
97 (PAS) is commonly used with TD for detection to satisfy its high time resolution demands. As an in-
98 situ technique, PAS eliminates the artifacts associated with filter-based methods (Weingartner et al.,
99 2003; Coen et al., 2010) and is often considered as the reference instrument for light absorption
100 coefficient determination (Arnott et al., 2003; Arnott et al., 2005). One limitation of the TD approach is
101 that a universal optimal operation temperature does not exist, leading to a chemical composition
102 dependent efficiency. If the temperature is too low, the coating cannot be fully removed, and charring
103 can occur if the TD temperature is too high, leading to biased results. For example, a SP-AMS study
104 in Toronto found that the efficiency of BC coating removal by TD decreased substantially for wildfire
105 influenced samples (Healy et al., 2015). Another issue is particle loss due to TD, which can be ~ 20%
106 and needs to be taken into account (Ueda et al., 2016). It's also worth noting that MAE_p by the TD
107 approach is different from the MAE_p at the emission source. First, the morphology of thermally
108 denuded BC particles (compact aggregates) is different from that of freshly emitted BC particles
109 (chain-like aggregates). Second, most of the coatings is removed with the TD denuded soot particles,
110 but freshly emitted soot particles usually come with a thin coating of OC formed from condensation
111 of OC vapors as the temperature drops from engine to the ambient air. As a result, the MAE_p by TD

112 approach is expected to be lower than the MAE_p of emission source. In this sense, the TD approach
113 may not be a perfect “time machine” to reverse the aging process for E_{abs} determination.

114 The second approach is the MAE ratio method, which is also stated in Eq. 3. The key to this method
115 is determining an appropriate MAE_p that can represent the MAE from primary emissions. One
116 approach is to adopt the reference MAE_p from the literature but it may fail to represent the actual
117 MAE_p at a specific sampling site, since MAE_p varies temporally and spatially. For example, MAE_p of
118 diesel soot was found to be 7.1 m²g⁻¹ at 532 nm (Adler et al., 2010). A much higher MAE_p (16 m²g⁻¹
119 at 530 nm) was observed from natural gas flaring (Weyant et al., 2016). MAE_p of biomass burning
120 (BB) samples is highly varied due to a wide range of fuel types and combustion conditions (Reid et
121 al., 2005; Roden et al., 2006). A range from 6.1 to 80.8 m²g⁻¹ was reported for BB MAE_p at 550 nm
122 (Pandey et al., 2016). Without the knowledge of source contributions, it is not feasible to derive a
123 representative MAE_p for E_{abs} estimation. The other commonly used approach is to determine MAE_p
124 from the dependency of MAE on the number fraction of coated soot particles measured by SP2 (Lan
125 et al., 2013). Since MAE (y axis) is positively correlated with the number fraction of coated soot
126 particles (x axis), MAE_p can be determined by extending the regression line to x=0. It is worth noting
127 that this approach provides only a rough approximation of E_{abs} since the parameter used here (coated
128 soot particles number fraction) ignores other main drivers of light absorption enhancement (e.g.
129 coating thickness). As a result, this approach is only valid for a period of measurements, for which
130 coating thickness is relatively constant and the MAE variations are dominated by coated soot particles
131 number fraction.

132 However, the high cost of the TD-PAS system and SP2 limit the field measurement of E_{abs} around
133 the world. In addition, long-term E_{abs} measurements by a TD-PAS system and SP2 are not easily
134 achieved and rarely reported. On the other hand, an Aethalometer and RT-ECOC analyzer can be
135 effectively deployed for long term measurements and E_{abs} estimation, at a relatively lower cost. In this
136 study, based on one year of hourly MAE measurements (with the field carbon analyzer and
137 Aethalometer) at a suburban site in the Pearl River Delta (PRD) region of China, quantification of
138 MAE_p is demonstrated by a novel statistical approach, the Minimum R squared method (MRS) (Wu
139 and Yu, 2016). The aim of this study is to demonstrate the capability of E_{abs} estimation using a year-

140 long dataset from cost-effective instrumentation. The seasonal variability of MAE, AAE and E_{abs} in
141 the PRD region are characterized and their dependency on air mass origin and biomass burning are
142 discussed. Abbreviations used in this study are summarized in Table 1 for a quick lookup.

143 **2 Ambient measurements**

144 Sampling was conducted from Feb 2012 to Jan 2013 at the suburban Nancun (NC) site (23° 0'11.82"N,
145 113°21'18.04"E). NC, situated on the top of the highest peak (141 m ASL) in Guangzhou's Panyu
146 district, is located at the geographic center of the PRD region, making it a representative location for
147 average atmospheric mixing characteristics of city clusters in the PRD region. Light absorption
148 measurements were performed by a 7- λ Aethalometer (AE-31, Magee Scientific Company, Berkeley,
149 CA, USA). The Aethalometer was equipped with a 2.5 μ m cyclone with a sampling flow rate of 4 L
150 min^{-1} . Weingartner's algorithm (Weingartner et al., 2003) was adopted to correct the sampling artifacts
151 (aerosol loading, filter matrix and scattering effect) rooted in filter based method. A customized
152 Aethalometer data processing program (Wu, 2017a) with graphical user interface was developed to
153 perform data correction and detailed descriptions can be found in the SI (The program is available
154 from <https://sites.google.com/site/wuchengust>). Details of the Aethalometer setup and data correction
155 can be found in our previous paper (Wu et al., 2013).

156 EC mass concentrations were determined by a real time ECOC analyzer (Model RT-4, Sunset
157 Laboratory Inc., Tigard, Oregon, USA). The sunset carbon analyzer was sampling on hourly cycles at
158 a flow rate of 8 Lmin^{-1} with a $\text{PM}_{2.5}$ sharp-cut cyclone inlet. For each measurement hour, the first
159 45min were for sample collection and the remaining 15 min for thermal-optical analysis. OC is
160 volatilized first by step-wise temperature ramping in an oxygen-free atmosphere while in the second
161 stage EC is combusted in the presence of oxygen. Laser transmittance is applied to correct the charring
162 artifact during the OC stage.

163 Considering a measurement precision of 5% for the Aethalometer (Hansen, 2005) and 24% for the
164 RT-ECOC analyzer (Bauer et al., 2009), the propagated relative precision of E_{abs} ($E_{abs,Unc}$) is 35%
165 following Eq. S1&S2 in the SI. It should be noted that $E_{abs,Unc}$ is mainly attributed to the

166 measurement precision of EC by the RT-ECOC analyzer. Since the measurement precision of the RT-
167 ECOC analyzer estimated by Bauer et al. (2009) is obtained from field measurement at an environment
168 (EC below $1 \mu\text{g m}^{-3}$) where EC is much lower than the present study (annual average EC 2.66 ± 2.27
169 $\mu\text{g m}^{-3}$), the $E_{abs,Unc}$ of 35% should be considered as an upper limit for the present study.

170 Light scattering was measured by an integrating nephelometer (Aurora-1000, Ecotech, Melbourne,
171 Australia). Water soluble ions were measured by MARGA (The instrument for Measuring AeRosols
172 and GAses)(ten Brink et al., 2007). Both instruments are equipped with a $\text{PM}_{2.5}$ inlet to remove the
173 coarse particles.

174 **2.1 Uncertainties of MAE determination**

175 Two major uncertainties associated with the σ_{abs} and EC determination techniques should be taken
176 into account when comparing MAE across different studies. For the σ_{abs} determination technique,
177 photo-acoustic spectroscopy (PAS) is an in-situ technique free from filter based artifacts, but its
178 application is limited by its high cost. The filter based optical transmittance method (e.g., Aethalometer
179 and Multi Angle Absorption Photometer, MAAP) is the most widely used technique around the world,
180 but data correction is needed to minimize the bias from artifacts due to the loading effect, matrix effect
181 and scattering effect (Weingartner et al., 2003; Arnott et al., 2005; Schmid et al., 2006; Virkkula et al.,
182 2007; Coen et al., 2010; Drinovec et al., 2017; Saturno et al., 2017). Besides these artifacts, RH is also
183 a source of σ_{abs} measurement uncertainty. Elevated RH is not only a driving force of increased σ_{abs}
184 due to the hygroscopic growth of particles, but also a factor affecting ambient σ_{abs} measurements.
185 Previous studies found σ_{abs} by PAS exhibit a systematic decrease when $\text{RH} > 70\%$ (Arnott et al., 2003;
186 Kozlov et al., 2011). Water evaporation was found as the major cause for the biased PAS σ_{abs}
187 measurements under high RH (Raspet et al., 2003; Lewis et al., 2009b; Langridge et al., 2013). Filter-
188 based measurements are also affected under high RH conditions. For example, Arnott et al. (2003)
189 observed erratic responses by particle soot absorption photometer (PSAP) as RH changed. The main
190 reason is traced to the hydrophilic cellulose membrane, which serves to reinforce the quartz filter used
191 in PSAP. The fibers can swell and shrink as RH changes, causing unwanted light attenuation signal.
192 The PTFE-coated glass-fiber tape has become available since 2012 for the recent model of
193 Aethalometer to minimize the RH interference (Magee-Scientific, 2017). A study by Schmid et al.
194 (2006) reported dependency of PSAP σ_{abs} on RH, but found negligible effect of RH on Aethalometer
195 performance. It is also worth noting that RH in the Aethalometer optical chamber may be lower than

196 the ambient RH due to the slightly elevated temperature inside the instrument. The magnitude of RH
197 difference was found similar between different instruments: 20% for the Aethalometer (Schmid et al.,
198 2006) and 15% for the nephelometer (Guyon et al., 2004). The RH in the Aethalometer optical chamber
199 was not measured in this study. We expected its level to be slightly lower than the ambient RH. Cappa
200 et al. (2008) found σ_{abs} measurements by PSAP and PAS maintained a high linearity even under high
201 RH conditions (65-91%). Inter-comparison studies demonstrated that with proper corrections,
202 Aethalometer σ_{abs} measurements agree well with those by PAS (Ajtai et al., 2011). During the inter-
203 comparison study of an Aethalometer (AE-16) and a PAS in Guangzhou (Wu et al., 2009), good
204 correlation was found ($R^2=0.96$) as shown in Figure S1. These comparison results imply that the
205 Aethalometer results are linearly correlated with PAS measurements and RH has a limited interference
206 on Aethalometer measurements. In our study, careful corrective measures (Wu et al., 2013) are
207 conducted for the Aethalometer σ_{abs} data treatment to minimize these artifacts. But such artifacts
208 still cannot be fully eliminated.

209 For the EC determination, different thermal optical analysis (TOA) protocols can impact the
210 measurement variability and thus MAE. As shown in Table S1, MAE for the same samples at Fresno
211 varied from 6.1 to 9.3 $\text{m}^2 \text{g}^{-1}$, depending on which EC analysis protocol was applied (Chow et al.,
212 2009). Studies in the PRD found that discrepancies in measured EC by different analysis protocols
213 could be as large as a factor of 5 (Wu et al., 2012; Wu et al., 2016a), which adds to the uncertainty for
214 the MAE estimation. In addition, EC by TOA is also different from refractory BC (rBC) reported by
215 the laser induced incandescence (LII) technique (e.g. single particle soot photometer, SP2). For
216 example, two studies in Toronto (Knox et al., 2009; Chan et al., 2011) both used the PAS for σ_{abs}
217 measurement but different techniques for EC mass determination, resulting in very different MAE
218 results. LII instruments are usually calibrated with a commercially available surrogate (e.g. fullerene)
219 since direct calibration with ambient soot is not easy to achieve. Laborde et al. (2012) indicates that
220 the incandescence response of SP2 exhibits a dependency on soot type (15% between fullerene and
221 denuded diesel soot particles; 14% between biomass burning and denuded diesel soot particles). Due
222 to the absence of widely accepted reference materials for EC, the uncertainties in EC determination
223 will exist in the foreseeable future. All these uncertainties, including the uncertainty of rBC mass
224 determination by SP2, uncertainty of EC in TOA, the discrepancy between SP2 rBC and TOA EC and

225 the discrepancy of σ_{abs} between filter transmission and photo-acoustic methods, can contribute to the
226 differences in MAE listed in Table S1.

227 Systematic bias in MAE (e.g. overestimation of σ_{abs} and variability of EC mass by different
228 TOA protocols) discussed above have little effect on E_{abs} estimation by MRS. As shown in Eq. 3, E_{abs}
229 is the ratio of MAE_t to MAE_p or $\sigma_{abs,t}$ to $\sigma_{abs,p}$, thus most of the bias in EC mass or σ_{abs} is
230 cancelled out during the E_{abs} calculation. More details are discussed in section 4.1.

231 **3 Methodology**

232 **3.1 MAE_p estimation by MRS from the ambient data**

233 In this section, a new approach for MAE_p estimation is introduced for E_{abs} determination, which
234 requires the knowledge of differentiating $\sigma_{abs,p}$ and $\sigma_{abs,c}$ portions in $\sigma_{abs,t}$. The idea of
235 decoupling $\sigma_{abs,t}$ into $\sigma_{abs,p}$ and $\sigma_{abs,c}$ is conceptually similar to decoupling OC into primary OC
236 (POC) and secondary OC (SOC) in the EC tracer method as shown in Table 2. In the EC tracer method,
237 if $(OC/EC)_p$ is known, POC can be determined from OC (Turpin and Huntzicker, 1991). The role of
238 MAE_p here is similar to the role of $(OC/EC)_p$, the primary OC/EC ratio in the EC tracer method (a
239 comparison is given in Table 2). If MAE_p (average MAE from primary emission sources) is known,
240 E_{abs} can be obtained from the ratio of MAE_t/MAE_p (Eq. 3). Therefore, the key for E_{abs} estimation is to
241 derive an appropriate MAE_p. It is worth noting that MAE_p here does not represent MAE from a single
242 or specific primary emission source, instead it reflects an average and effective MAE that has taken
243 consideration of various primary emission sources. Thus, the MAE_p is conceptually analogous to
244 $(OC/EC)_p$ in the EC tracer method, in which the primary ratio reflects an overall ratio from primary
245 emission sources rather than from a single primary source.

246 The Minimum R squared method (MRS) explores the inherent independency between
247 pollutants from primary emissions (e.g., EC) and products associated with secondary formation
248 processes (e.g., SOC, $\sigma_{abs,c}$) to derive the primary ratios (e.g., $(OC/EC)_p$, MAE_p) in the EC tracer
249 method (Wu and Yu, 2016). When applying MRS for light absorption enhancement estimation, MRS
250 is used to explore the inherent independency between EC and $\sigma_{abs,c}$, which is gained during

251 atmospheric aging after emission. An example of MAE_p estimation by MRS is shown in Figure 1.
252 Firstly, the assumed MAE_p value is varied continuously in a reasonable range (0.01 to 50 $m^2 g^{-1}$ as
253 shown in Figure 1). Then at each hypothetical MAE_p , $\sigma_{abs,c}$ can be calculated by Eq. 6 (a combination
254 of Eq. 2&4) using EC and $\sigma_{abs,t}$ from ambient measurements.

$$255 \quad \sigma_{abs,c} = \sigma_{abs,t} - MAE_p \times EC \quad (6)$$

256 Accordingly, for each hypothetical MAE_p , a correlation coefficient value (R^2) of $\sigma_{abs,c}$ vs.
257 EC (i.e., $R^2(\sigma_{abs,c}, EC)$) can be obtained. The series of $R^2(\sigma_{abs,c}, EC)$ values (y axis) are then plotted
258 against the assumed MAE_p values (x axis) as shown by the red curve in Figure 1. The physical meaning
259 of this plot can be interpreted as follows. The $\sigma_{abs,p}$ is the fraction of light absorption owing to
260 primary emitted soot particles. As a result, $\sigma_{abs,p}$ is well correlated with EC mass. In contrast, the
261 $\sigma_{abs,c}$ is the fraction of light absorption gained by the lensing effect of the coating on particles after
262 emission. The variability of $\sigma_{abs,c}$ mainly depends on the coating thickness of the soot particles.
263 Consequently, $\sigma_{abs,c}$ is independent of EC mass. Since variations of EC and $\sigma_{abs,c}$ are independent,
264 the assumed MAE_p corresponding to the minimum $R^2(EC, \sigma_{abs,c})$ would then represent the most
265 statistically probable MAE_p of the tested dataset.

266 A computer program (Wu, 2017b) in Igor Pro (WaveMetrics, Inc. Lake Oswego, OR, USA)
267 was developed to facilitate MRS calculation with a user friendly graphical user interface. Another two
268 Igor Pro based computer programs Histbox (Wu, 2017c) and Scatter Plot (Wu, 2017d) are used for
269 generating histograms, box plots and scatter plots (with Deming regressions) presented in this study.
270 Detailed descriptions of these computer programs can be found in the SI and the computer programs
271 are available from <https://sites.google.com/site/wuchengust>.

272 **3.2 Mie simulation**

273 It can be informative to model a single soot particle using Mie theory (Bohren and Huffman,
274 1983) and understand the theoretical range and variability of the soot particle's optical properties.
275 Three types of mixing state are widely employed for parameterization: internal mixing, external
276 mixing and core-shell. To better represent the real situation (coating due to the aging process), a core-

277 shell model is considered in the Mie calculation (Figure S2), which is more realistic than a volume
278 mixture model (Bond et al., 2006). An aerosol optical closure study in the North China Plain (NCP)
279 found that the core-shell model can provide better performance than assuming purely internal mixing
280 and external mixing (Ma et al., 2012). A morphology study using Scanning Transmission X-ray
281 Microscopy found that core-shell is the dominating mixing state in ambient samples (Moffet et al.,
282 2016). It should be noted that the core-shell model assumption still has its own limitations. A single
283 particle soot photometer (SP2) study by Sedlacek et al. (2012) reported a negative lag time between
284 the scattering and incandescence signals in samples influenced by biomass burning, implying a near
285 surface location of soot relative to non-absorbing materials. Near surface type mixing of soot has also
286 been observed in Tokyo, but accounted for only 10% of total mixed soot containing particles (Moteki
287 et al., 2014). Considering the domination of core-shell type particles in the ambient environment, the
288 core-shell assumption in our optical model is sufficient to approximate the real situation.

289 As shown in Figure S2, fresh emitted soot particles are chain-like aggregates of small spheres
290 (30~50 nm). After the aging process, soot particles are coated with organic and inorganic materials.
291 Sufficient evidence has shown that the coating not only results in particle size growth, but also makes
292 the soot core become more compact due to its collapse (Alexander et al., 2008; Zhang et al., 2008;
293 Lewis et al., 2009a), especially under high RH conditions (Leung et al., 2017). A recent study by Pei
294 et al. (2017) shown that filling of void space within the agglomerate is the first step of the
295 morphological transformation of soot particles in atmospheric aging, leading to a spherical soot core.
296 Since the spherical like core and shell favor Mie simulation, both core and shell are considered as
297 spheres in the Mie calculation.

298 To investigate the spectrum properties of soot particles, 11 wavelengths (370, 405, 470, 520,
299 532, 550, 590, 660, 781, 880 and 950 nm) are considered in calculations to cover wavelengths in the
300 most frequently used absorption measurement instruments. A refractive index (RI) of $1.85 - 0.71i$ is
301 adopted for soot core (Bond and Bergstrom, 2006) and 1.55 for non-absorbing coating (clear shell) in
302 the Mie calculation for all wavelengths. Studies suggest a group of organic matter (OM), known as
303 Brown Carbon (BrC), can absorb solar radiation at UV wavelengths (Kirchstetter et al., 2004). Thus,
304 a BrC coating (brown shell) scenario is also considered in Mie simulation following the wavelength-

305 dependent RI suggested by Lack and Cappa (2010), which ranges from 1.55-0.059i (370 nm) to 1.55-
306 0.0005i (950 nm). A modeling study by Bond et al. (2006) indicates that absorption amplification is
307 not sensitive to the RI, thus the result below is not expected to be sensitive to the RI variability. Due
308 to the spherical assumption of the BC core, a constant particle density is adopted for simplicity instead
309 of size dependent particle density. But it is worth noting that in reality, the effective density of soot
310 varies with particle size due to the morphology change during particle aging (Tavakoli and Olfert,
311 2014; Dastanpour et al., 2017). Both core diameters (D_{core}) and shell diameters (D_{shell}) are constrained
312 in the range of 10 ~ 3000 nm in the model simulations. The Mie calculations are implemented with a
313 customized program (Wu, 2017e) written in Igro Pro (WaveMetrics, Inc. Lake Oswego, OR, USA)
314 and it is available from <https://sites.google.com/site/wuchengust>. It should be noted that the core-shell
315 type mixing state of particles is still rare in 3D atmospheric models like WRF-Chem (Matsui et al.,
316 2013; Nordmann et al., 2014) due to computational cost limitation.

317 **3.2.1 Mie modeled absorption angstrom exponent (AAE)**

318 Absorption Angstrom Exponent (AAE) is a widely used parameter that describes the
319 wavelength dependence of aerosol light absorption (Moosmuller et al., 2011), which can be written
320 explicitly as

$$321 \quad AAE(\lambda_1, \lambda_2) = -\frac{\ln(\sigma_{abs,\lambda_1}) - \ln(\sigma_{abs,\lambda_2})}{\ln(\lambda_1) - \ln(\lambda_2)} \quad (7)$$

322 It is well known that ambient soot particles exhibit an AAE close to unity (Bond, 2001).
323 Modeled variability in $AAE_{470-660}$ of bare soot particles is shown in Figure S3. For soot particles with
324 $D_{\text{core}} < 200$ nm, $AAE_{470-660}$ is very close to 1 and decreases significantly for particles with $D_{\text{core}} > 200$
325 nm. Considering a typical D_{core} of fresh emitted soot particles smaller than 200 nm (Rose et al., 2006;
326 China et al., 2013), the model results confirm the frequently observed AAE close to 1 from ambient
327 measurements (Kirchstetter et al., 2004). Modeled variability in $AAE_{470-660}$ of soot particles coated by
328 non-absorbing substances (clear shell) and weakly absorbing materials (brown shell) is shown in
329 Figure 2. Elevated AAE to ~2 is observed in the clear shell scenario (Figure 2a and 3b) for the most
330 probable soot core particle sizes (<200 nm), which agrees well with a previous model study (Lack and

331 Cappa, 2010), implying that elevated AAE cannot be exclusively attributed to mixing with BrC. AAE
332 elevation is more pronounced in the brown shell scenario. For soot particles with $D_{\text{core}} < 200$ nm, brown
333 shell $AAE_{470-660}$ can easily reach 3 for a coating of $D_{\text{shell}}/D_{\text{core}}=3$ (Figure 2c and 2d). These high AAE
334 results are consistent with the previous model study (Lack and Cappa, 2010) and could partially
335 explain the high AAE observed in measurement studies (Kirchstetter et al., 2004; Hoffer et al., 2006),
336 since the presence of externally mixed BrC particles also contribute to the wavelength dependent light
337 absorption.

338 **3.2.2 Mie modeled single scattering albedo (SSA)**

339 Variability in modeled SSA_{525} of soot particles coated by non-absorbing substances and weakly
340 absorbing materials (e.g. BrC) is shown in Figure S4. For particles with $D_{\text{core}} < 200$ nm and $D_{\text{shell}}/D_{\text{core}}$
341 < 3 , the SSA increases gradually (up to ~ 0.9) with a thicker coating and behaves similarly between
342 clear shell and brown shell scenarios.

343 **3.2.3 Mie modeled mass absorption efficiency (MAE)**

344 MAE is a useful indicator for soot mixing state. Variability in MAE of bare soot particles as a
345 function of particle size at a wavelength of 550 nm is illustrated in Figure S5. The magnitude of MAE
346 is sensitive to the soot density assumption, especially for particles < 200 nm (Figure S5), but the overall
347 trend of particle size dependency is similar between different density scenarios. MAE peaks at a
348 particle size of 200 nm and decreases dramatically for larger particles. In our MAE calculation, a soot
349 density of 1.9 g cm^{-3} is adopted, as suggested by Bond and Bergstrom (2006). The purpose of adopting
350 constant density is to simplify the MAE calculation. It should be noted that the effective density of
351 soot core is highly variable in ambient environments. For example, a study in Beijing (Zhang et al.,
352 2016b) found a value of 1.2 g cm^{-3} . A recent chamber study found the effective density of soot can
353 evolve from 0.43 to 1.45 g cm^{-3} during aging as coated by m-Xylene oxidation products (Guo et al.,
354 2016). A study by a single-particle aerosol mass spectrometer in Guangzhou found the effective
355 density of soot increased with particle size in the range of 400 to 1600 nm (Zhang et al., 2016a). The
356 MAE of coated particles from different core/shell diameter combinations are shown in Figure S6. For

357 thickly coated particles, the MAE in the clear shell scenario varied as $D_{\text{shell}}/D_{\text{core}}$ increased, but the
358 MAE of brown shell scenario increased quasi-monotonously with $D_{\text{shell}}/D_{\text{core}}$.

359 **3.2.4 Mie modeled light absorption enhancement factor (E_{abs})**

360 E_{abs} is a better indicator for soot mixing state than MAE since it does not rely on the soot density
361 assumption and is more suitable for comparing Mie simulations with ambient measurements. Modeled
362 variability in E_{abs} of soot particles coated by non-absorbing substances and weakly absorbing materials
363 (e.g. BrC) is shown in Figure 3a and 3c respectively. E_{abs} is not only sensitive to the core/shell diameter
364 combination, but also behaves very differently on the clear and brown shell assumptions. For the clear
365 shell scenario, when $D_{\text{coat}}/D_{\text{core}} < 2$, E_{abs} does not exceed 2 for particles with different soot core sizes,
366 but for the same $D_{\text{coat}}/D_{\text{core}}$, a larger soot core size yields a higher E_{abs} (Figure 3b, cross-sections of
367 Figure 3a). If $D_{\text{coat}}/D_{\text{core}} > 2$, E_{abs} could be 3 to 5 for particles with a soot core smaller than 200 nm, but
368 for particles with a soot core larger than 200 nm, the E_{abs} is limited to ~ 2 as shown in Figure 3b. For
369 the brown shell scenario, E_{abs} increased quasi-monotonically with $D_{\text{coat}}/D_{\text{core}}$, and this trend is similar
370 for different soot core sizes (Figure 3d). The main reason behind is that in the brown shell scenario,
371 both lensing effect and BrC absorption contribute to E_{abs} . As shown in Figure S7, the BrC absorption
372 contribution to total E_{abs} strongly depends on coating thickness and is insensitive to soot core diameters.
373 When the coating is relatively thin (< 5 nm for $\lambda@370$ nm, < 15 nm for $\lambda@550$ nm and < 40 nm for
374 $\lambda@880$ nm), BrC absorption contribution to the total E_{abs} is less than 20%. As the coating increases to
375 a certain level (~ 15 nm for $\lambda@370$ nm, ~ 35 nm for $\lambda@550$ nm and ~ 90 nm for $\lambda@880$ nm), BrC
376 absorption contribution is comparable to the lensing effect contribution, each contributing $\sim 50\%$ to
377 the total E_{abs} . When the BrC coating is sufficiently thick (> 30 nm for $\lambda@370$ nm, > 90 nm for $\lambda@550$
378 nm and > 110 nm for $\lambda@880$ nm), BrC absorption dominates the E_{abs} contribution. As a result, if BrC
379 coating is indeed present in ambient samples, a strong wavelength dependent E_{abs} could be observed,
380 since a BrC coating of 30 nm would be enough to induce a large amount of detectable E_{abs} in the UV
381 range. Another major difference between the clear and brown shell scenarios is that, for thickly coated
382 particles (e.g. $D_{\text{coat}}/D_{\text{core}} > 2$), the brown shell can yield a much higher E_{abs} than the clear shell.

383 Both primary soot size distribution and coating thickness can affect the absorption
384 enhancement of ambient BC particles. Ambient measurements by LII found soot particle number and
385 mass modes peaking at 110 nm and 220 nm, respectively, in the PRD (Huang et al., 2011). A study in
386 Shanghai found similar results (70 nm for number concentrations and 200 nm for mass
387 concentrations)(Gong et al., 2016). Considering that the LII technique is specific for BC mass
388 determination which is independent of BC mixing state, the size distribution reported by LII can
389 represent the size distribution of the BC core. A study using a Micro Orifice Uniform Deposit Impactor
390 (MOUDI) found a EC mass size distribution in the PRD exhibiting three modes peaking at ~300, ~900
391 and ~5000 nm (Yu et al., 2010), implying a substantial coating of BC particles, and a diameter
392 amplification of 3. BC sizing by LII is based on volume equivalent diameter (VED), while MOUDI is
393 based on aerodynamic diameter. As a result, these two techniques do not necessarily yield similar sizes,
394 even for the bare soot particles. The conversion between these two types of diameters involves the
395 knowledge of particle density and morphology (drag force). A recent closure study on BC mixing state
396 in the PRD region suggests σ_{abs} is dominated by coated soot particles in the range of 300~400 nm
397 (Tan et al., 2016). Considering the dominant BC core distribution measured by SP2 (110 nm), the
398 upper limit of E_{abs} in the PRD is roughly estimated as ~2 for the clear shell scenario (Figure 3b).

399 **4 Results and discussions**

400 **4.1 Annual measurement statistics**

401 The frequency distribution (log-normal) of σ_{abs550} is shown in Figure 4a, with an annual average (± 1
402 S.D.) of $42.65 \pm 30.78 \text{ Mm}^{-1}$. A log-normal distribution is also found in the EC mass concentration
403 (Figure 4b), with an annual average of $2.66 \pm 2.27 \mu\text{g m}^{-3}$. Figure 4c demonstrates the yearlong
404 frequency distribution of MAE_{550} at the NC site. The annual average MAE_{550} is $18.75 \pm 6.16 \text{ m}^2 \text{ g}^{-1}$ and
405 the peak (± 1 S.D.) of the lognormal fit is $15.70 \pm 0.22 \text{ m}^2 \text{ g}^{-1}$. A good correlation is observed between
406 σ_{abs} and EC mass ($R^2=0.92$) as shown in Figure 4d, and the color coding indicates a MAE dependency
407 on RH, which agrees with a study in Xi'an(Wu et al., 2016b). Annual average $AAE_{470-660}$ is 1.09 ± 0.13
408 (Figure S8a), indicating that soot is the dominant absorbing substance in the PRD and the brown shell

409 scenario shown in the Mie simulation is unlikely to be important. Annual mean SSA_{525} is 0.86 ± 0.05
410 (Figure S8c), similar to previous studies in the PRD (Jung et al., 2009; Wu et al., 2009). For
411 comparison purpose, MAE measured at original wavelength and MAE scaled to 550 nm following the
412 λ^{-1} assumption are both shown in Table S1. The MAE comparisons discussed below are MAE at 550
413 nm. MAE_{550} by previous studies at various locations was found to cover a wide range, from 5.9 to 61.6
414 $m^2 g^{-1}$. Annual average observed MAE_{550} at NC ($18.75 m^2 g^{-1}$) is higher than many studies shown in
415 Figure 5, e.g., Shenzhen (Lan et al., 2013), Beijing (Yang et al., 2009), Mexico city (Doran et al., 2007)
416 and Fresno (Chow et al., 2009).

417 As shown in Figure 1, the annual average $MAE_{p,550}$ estimated by MRS is $13 m^2 g^{-1}$. MAE_p by
418 MRS represents the MAE_p at the emission source, which is different from the MAE_p by the TD
419 approach for two reasons. First, the morphology of thermally denuded BC particles (compact
420 aggregates) is different from that of freshly emitted BC particles (chain-like aggregates). Second, most
421 of the coatings are removed for TD denuded soot particles, but freshly emitted soot particles usually
422 come with a thin coating of OC formed from condensation of OC vapors as the temperature drops from
423 the flame to the ambient air. As a result, the MRS-derived MAE_p is expected to be higher than the
424 MAE_p by the TD approach. The estimated $MAE_{p,550}$ is higher than a previous study in Guangzhou
425 ($7.44 m^2 g^{-1}$) (Andreae et al., 2008), but comparable to Xi'an ($11.34 m^2 g^{-1}$) (Wang et al., 2014) and
426 Toronto ($9.53\sim 12.57 m^2 g^{-1}$) (Knox et al., 2009). The annual average E_{abs550} by MRS following Eq. 3
427 is estimated to be 1.50 ± 0.48 (mean ± 1 S.D.).

428 As mentioned in section 1, the definition of MAE_p by the TD approach is different from the
429 MAE_p of emission source. The TD MAE_p is expected to be slightly lower than the MAE_p of emission
430 source. Therefore, the corresponding E_{abs} are slightly different and it should be cautioned when
431 comparing MRS-derived E_{abs} with E_{abs} by the TD approach and Mie simulations. The E_{abs} could vary
432 by location, depending on the coating thickness and size distribution of the primary aerosols. After
433 undergoing atmospheric aging, the E_{abs} can be increased during transport from emission source to rural
434 areas. The magnitude of the E_{abs} found at the NC site is comparable to other locations such as Boulder
435 (Lack et al., 2012a) (1.38), London (Liu et al., 2015) (1.4), Shenzhen (Lan et al., 2013) (1.3), Yuncheng
436 (Cui et al., 2016b) (2.25), Jinan (Chen et al., 2017) (2.07) and Nanjing (Cui et al., 2016a) (1.6) and is

437 higher than studies in California (Cappa et al., 2012) (1.06), as listed in Table 3. Spectrum E_{abs} are
438 calculated from 370 to 950 nm as shown in Figure S9. E_{abs} in the PRD exhibits a weak wavelength
439 dependence, with slightly higher E_{abs} at the shorter wavelength (e.g. $E_{\text{abs}370} = 1.55 \pm 0.48$) and is
440 relatively lower in the IR range (e.g. $E_{\text{abs}950} = 1.49 \pm 0.49$).

441 **4.2 Monthly characteristics of MAE, AAE and SSA**

442 Monthly variations of MAE_{550} at the NC site are shown in Figure 6a and Table S2, revealing distinct
443 patterns of higher MAE_{550} in summer and lower in winter. On the other hand, $\text{AAE}_{470-660}$ is lower in
444 summer and higher in winter (Figure 6b and Table S3). Monthly SSA_{525} varied from 0.83 to 0.90
445 without a clear seasonal pattern, as shown in Figure S10 and Table S4. $\text{MAE}_{\text{p},550}$ estimation for
446 individual months is shown in Figure 6a (the purple line) and monthly $E_{\text{abs}550}$ is calculated accordingly
447 following Eq. 3 (Figure 6c). $E_{\text{abs}550}$ shows clear seasonal variations, with higher values from April to
448 August (1.52~1.97 as shown in Table S5) and relatively lower values from September to March
449 (1.24~1.49). The highest enhancement is found in August (1.97). Factors affecting variation of $E_{\text{abs}550}$
450 are discussed in the following sections, including air mass origin and biomass burning.

451 **4.3 The effect of air mass origin**

452 It's of interest to understand the seasonal variations of optical properties in the PRD. Hourly backward
453 trajectories for the past 72 hours were calculated using NOAA's HYSPLIT (Hybrid Single Particle
454 Lagrangian Integrated Trajectory, version 4) model (Draxier and Hess, 1998) from Feb 2012 to Jan
455 2013 as shown in Figure S11. Cluster analysis was conducted using MeteoInfo (Wang, 2014). By
456 examining the total spatial variance (TSV), the number of clusters was determined to be four as shown
457 in Figure S12. Cluster 1 (C1) represents continental air masses from the north, accounting for 44.4%
458 of total trajectories. C2 (22.8%) represents marine air masses coming from the South China Sea. C3
459 represents air masses from the east (Taiwan island). C4 (15.8%) represents transitional air masses
460 coming from the east coastline of China. As shown in Figure 7, $E_{\text{abs}550}$ from C2 (1.78) is higher than
461 other clusters (1.30 – 1.42). Further Wilcoxon-Mann-Whitney tests show that $E_{\text{abs}550}$ from C2 is
462 significantly higher than $E_{\text{abs}550}$ from C1, C3 and C4 (Figure S13), implying that particles from the

463 South China Sea cluster is likely more aged than other clusters. Air mass origin in the PRD is
464 dominated by C2 from Apr to Aug (Figure S14a) as a result of the South China Sea monsoon in the
465 rainy season. In contrast, the dry season is ruled by continental air masses from the north (C1) due to
466 the influence of the northeast monsoon. $E_{\text{abs}550}$ from C2 varied from 1.67 to 2.19, but was always
467 higher than $E_{\text{abs}550}$ from C1 and C3 during the rainy season (Figure S14b). As a result, the domination
468 of aged air mass from the vast ocean is one of the reasons for the much higher $E_{\text{abs}550}$ found in the rainy
469 season.

470 **4.4 The effect of biomass burning**

471 Biomass burning (BB) and vehicular emission are the two major sources of soot particles. BC
472 from biomass burning emission, depending on the fuel type and burning condition, may have a higher
473 OC/EC ratio and a thicker coating, resulting in a higher MAE than vehicular emission (Shen et al.,
474 2013; Cheng et al., 2016). In this study, the influence of BB on optical properties is investigated using
475 the K^+/EC ratio as a BB indicator. As shown in Figure 8, MAE_{550} is positively correlated with the
476 K^+/EC ratio, which exhibits a clear seasonal pattern that is higher in the rainy season and lower in the
477 dry season (Figure S15a). Southeast Asia has the highest fire emission density globally due to the high
478 biofuel consumption along with frequent fire activity in this region (Aouizerats et al., 2015), making
479 Southeast Asia a large contributor to BC emissions (Jason Blake, 2014). During the rainy season when
480 oceanic wind prevails, BC from BB emission in Southeast Asia can reach PRD through long range
481 transport (LRT), resulting in an elevated K^+/EC ratio and MAE_{550} . The Deming regression intercept
482 (11.89) in Figure 8 represents the MAE without the BB effect. This non-BB MAE_{550} ($11.89 \text{ m}^2 \text{ g}^{-1}$) is
483 only slightly lower than $\text{MAE}_{\text{p},550}$ ($13 \text{ m}^2 \text{ g}^{-1}$) obtained in section 4.3, implying that a large fraction of
484 $\text{MAE}_{\text{p},550}$ could not be explained by the BB source. Additional evidence was obtained through
485 examining regression relationships of $\text{MAE}_{\text{p},550}$ with K^+/EC month-by-month (Figure S15b).
486 Correlation of monthly $\text{MAE}_{\text{p},550}$ vs. K^+/EC ratio yield a R^2 of 0.23 (Figure S15c). In contrast, a much
487 higher correlation ($R^2=0.58$) was observed (Figure S15d) between $\text{MAE}_{\text{p},550}$ and non-BB MAE_{550} (i.e.,
488 K^+/EC intercepts from Figure S15b). These results imply that BB is one of the contributors to the
489 $\text{MAE}_{\text{p},550}$ variations, but unlikely the dominating one.

490 Many studies have found that BB influenced samples exhibit elevated AAE due to the presence
491 of wavelength dependent light absorbing substances like BrC and HUmic-Like Substances (HULIS)
492 (Kirchstetter et al., 2004; Hoffer et al., 2006; Sandradewi et al., 2008; Herich et al., 2011; Pokhrel et
493 al., 2017). It is of interest to investigate whether elevated AAE observed in the PRD during the dry
494 season is associated with BB influence. As shown in Figure S16, $AAE_{370-470}$ and $AAE_{470-660}$ did not
495 correlate with the BB indicator, K^+/EC ratio. These results suggest that the elevated AAE observed in
496 the PRD wintertime is unlikely to be dominated by the BB effect. Beside the independency between
497 $AAE_{470-660}$ and K^+/EC ratio, the measured $AAE_{470-660}$ range also implies that BB is not the major
498 driving force of $AAE_{470-660}$ variations. The limited light absorption contribution from BrC in RPD
499 region is observed in a recent study (Yuan et al., 2016) , which suggest an upper limit of BrC
500 contribution of 10% at 405 nm in the winter time using the AAE approach. As discussed in our Mie
501 simulation (section 3.1) and a previous study (Lack and Cappa, 2010), coating of non-absorbing
502 materials onto soot particles can increase AAE up to 2. Since the monthly average $AAE_{470-660}$ in
503 wintertime did not exceed 1.2 (Table S3), the variations of $AAE_{470-660}$ in the PRD are more likely
504 associated with coatings rather than the contribution of BrC. The results also imply that attempts on
505 BrC absorption attribution for the PRD dataset presented in this study could be risky, considering that
506 elevation of AAE is actually dominated by coating (Lack and Langridge, 2013).

507 **4.5 Implications for mixing state**

508 Quantitative direct measurements of BC mixing state and coating thickness are still challenging.
509 SP2 can estimate the coating thickness using a lag-time approach or a Mie calculation approach can
510 be employed, but both methods have a limited range in coating thickness and uncertainties arise from
511 the assumptions made during the retrieval. For example, recent studies found that the mass equivalent
512 diameter of soot core measured by SP2 could be underestimated due to density assumptions (Zhang et
513 al., 2016b). Although size distribution measurement is not available in this study, clues of mixing state
514 still can be derived from bulk measurements of optical properties. As discussed in section 4.4.1,
515 elevated E_{abs550} observed in the rainy season is associated with aged air masses from a marine origin.
516 To probe the possible mixing state difference between dry and rainy season, E_{abs550} , SSA_{525} and

517 AAE₄₇₀₋₆₆₀ are used to narrow down the possible core-shell size range as shown in Figure S17. Monthly
 518 averages with one standard deviation of AAE₄₇₀₋₆₆₀, SSA₅₂₅ and E_{abs550} are used as constraints to extract
 519 the intersecting core-shell size range from Figure 2a, Figure S4 and Figure 3a. January and August
 520 data are used to represent two different scenarios: elevated AAE₄₇₀₋₆₆₀ (1.19±0.11) with lower E_{abs550}
 521 (1.31±0.32) in dry season and low AAE₄₇₀₋₆₆₀ (1.04±0.09) with elevated E_{abs550} (1.97±0.71) in rainy
 522 season. The results show that January and August have a very different core-shell size range: in
 523 January, the core and shell range are 100 ~ 160 nm and 120 ~ 250 nm, respectively; in August, the
 524 core and shell range are 120 ~ 165 nm and 170 ~ 430 nm, respectively. This confirms again that the
 525 soot particles in the rainy season are likely to have a thicker coating than in the dry season.

526 **5 Caveats of the MRS method in its applications to ambient data**

527 **5.1 Impact of measurement biases**

528 It should be noted that the E_{abs} estimation approach is insensitive to the **systematic** MAE bias
 529 (e.g. **systematic** overestimation of σ_{abs} and variability of EC mass by different TOA protocols)
 530 discussed in section 2.1, because **systematic** bias in EC mass or σ_{abs} is cancelled out in the E_{abs}
 531 calculation (Eq. 3), since E_{abs} is the ratio of $\sigma_{abs,t}$ to $\sigma_{abs,p}$. To investigate the performance of the
 532 MRS approach in response to systematic bias in EC and σ_{abs} , two simple tests are conducted as shown
 533 in Figures S18 and S19 by adding systematic biases to the original data. **The one-year measurement**
 534 **data of σ_{abs550} and EC are used as original data.** Test A represents a situation when σ_{abs} is
 535 overestimated and EC is underestimated. The biased data are marked as σ'_{abs550} and EC' respectively,
 536 as shown below:

$$537 \quad \sigma'_{abs550} = \sigma_{abs550} \times 2 \quad (8)$$

$$538 \quad EC' = EC \times 0.7 \quad (9)$$

539 As a result, the average MAE₅₅₀ changed from 18.75 to 53.58 m² g⁻¹ and MAE_p changed from 13 to 37
 540 m² g⁻¹ (Figure S18). However, E_{abs} by ratio of averages remain the same (1.44).

541 In Test B, EC by different TOA protocols are compared to investigate the effect of different EC
 542 determination approaches while σ_{abs550} remains unchanged. EC by IMPROVE TOR protocol is
 543 calculated from NIOSH TOT EC following an empirical formula for suburban sites derived from a 3-
 544 year OCEC dataset in PRD (Wu et al., 2016a):

545
$$EC_{IMP_TOR} = 2.63 \times EC_{NSH_TOT} + 0.05 \quad (10)$$

546 As shown in Figure S19, MAE₅₅₀ changed from 18.75 to 7.02 m² g⁻¹ and MAE_p changed from 13 to 5
 547 m² g⁻¹, but E_{abs} remain almost the same (1.40). Result of Test B implies that although EC is
 548 operationally defined, the discrepancy of EC between TOA protocols did not weaken the role of EC
 549 serving as a tracer for primary emissions in MRS application. These examples demonstrate that
 550 systematic biases in σ_{abs550} and EC have no effects on E_{abs} estimation by the MRS approach.

551 Study by Cheng et al. (2016) found two distinct types of biomass smoke behave differently on the
 552 biases of filter based σ_{abs} measurement. The bias in the first type can be explained by a nearly
 553 constant correction factor, which is similar to the situation discussed in Test A. The bias in the second
 554 type shows an apparent OC/EC dependence. Test C is carried out to investigate this situation, i.e.,
 555 examining the impact of sample-dependent bias as a function of E_{abs}. Unlike the proportional bias in
 556 Test A and B that is the same for all data points, the bias in Test C depends on the E_{abs550} of individual
 557 samples, which are parametrized by Eqs. (11) and (12).

558
$$\sigma'_{abs550} = \sigma_{abs550} + \sigma_{abs550} \times (k \times E_{abs550} - k) \quad (11)$$

559
$$EC' = EC - EC \times (k \times E_{abs550} - k) \quad (12)$$

560 As shown in Eqs. (11) and (12), the positive bias of σ_{abs550} and negative bias of EC are proportional
 561 to E_{abs550}. The magnitude of E_{abs550}-dependent bias is regulated by the factor k. Since σ'_{abs550} and EC'
 562 are biased in different directions, resulting a further amplification in MAE biases, which could be
 563 considered as the extreme case. As shown in Figure S20, for k=10% (corresponding to a bias of 10%
 564 when E_{abs}=2), the bias of MRS-derived E_{abs} is very small (1%). For k=20%, the MRS-derived E_{abs}
 565 changes from 1.44 to 1.66, leading to a bias of 15%. These results imply that if the measurement bias
 566 follows the same form as demonstrated in Test C, the bias is not negligible but still acceptable. If the
 567 impact only affects σ_{abs} or EC rather than impacting both, the bias is expected to be smaller than the
 568 estimation shown in Test C.

569 It should be noted that the parameterization scheme shown in Eqs. (11) and (12) is only for
 570 demonstration purpose from a conceptual perspective and it does not necessarily represent the real-
 571 world measurements. There is a lack of quantitative understanding of this impact. For example, Lee et

572 al. (2007) used artificially fabricated EC samples with OC coatings to evaluate the impact of coating
573 on OC/EC analysis. Biases were observed, nevertheless, the results were linearly correlated with the
574 true OC and EC values with a high R^2 (>0.9), implying that the biases in that specific study were
575 dominated by systematic biases rather than the coating-dependent bias. Further studies are needed to
576 better characterize and parameterize this impact if filter-based techniques are used for σ_{abs} and EC
577 determination in the MRS approach.

578 **5.2 Impact of semi-volatile organic carbon**

579 Light absorption contribution due to semi-volatile organic carbon (SVOC) from wood combustion was
580 reported to be negligible in the visible range and around 10-20% at 360 nm (Chen and Bond, 2010).
581 On the other hand, OCEC analysis can be affected by SVOC (Subramanian et al., 2004), either by
582 positive artifacts through adsorption of SVOC onto quartz filters, or by negative artifacts through
583 evaporation of SVOC due to the gas-particle re-equilibrium down stream of VOC denuder. Positive
584 artifacts can be minimized by the installation of a VOC denuder which is widely adopted in RT-OCEC
585 measurements (Bae et al., 2004; Bauer et al., 2009). A typical negative artifact of 10% is expected and
586 can be corrected by backup filters (Subramanian et al., 2004). There was evidence to show that SVOC
587 could affect OC/EC split in thermal/optical analysis (Cheng et al., 2009). However, the bias in EC
588 caused by the OC/EC split drift due to SVOC is systematic, making it falls into the scenario discussed
589 in Test B. As a result, the impact from SVOC on E_{abs} estimation by MRS is expected to be small.

590 **5.3 Impact of mineral dust**

591 The presence of mineral dust (MD) could affect both σ_{abs} and EC determination. If MD is externally
592 mixed with soot particles, the light absorption from MD could be miscounted as σ_{abs} enhancement,
593 leading to the overestimation of E_{abs} . If the light absorption signal from MD is sufficiently strong
594 (e.g. $AAE > 2$), σ_{abs} by MD and BC can be separated by the AAE approach suggested by Fialho et al.
595 (2005). Additionally, the presence of substantial MD in samples has several impacts on the EC
596 determination by thermal analysis. First, if the samples are not pre-treated with acid, the carbonated
597 carbon could be misidentified as EC, resulting over-estimation of EC (Chow et al., 1993). The acid

598 treatment is only available for off-line OC/EC analysis and not yet practical for the RT-OCEC analyzer.
599 Second, metal oxides in MD can lead to premature EC oxidation in the helium stage of OC/EC analysis,
600 leading to underestimation of EC (Wang et al., 2010; Bladt et al., 2012). The lack of a parameterization
601 scheme for correcting the EC loss due to MD makes it improper to use the biased EC as a primary
602 tracer. For these reasons, E_{abs} estimation by MRS is not recommended for samples strongly influenced
603 by MD.

604 5.4 Impact of BrC

605 The data in this study is dominated by BC absorption that did not show much influence from
606 BrC. However, extra care should be taken if the samples exhibit substantial BrC signature (e.g.
607 $AAE > 2$). Such situations are equivalent to the two-source scenarios discussed in our previous paper
608 on the MRS method (Wu and Yu, 2016) and the major findings are described below. Two types of
609 two-source scenarios are considered: two correlated primary sources (scenario A) and two independent
610 primary sources (scenario B). In scenario A in which both BC and primary BrC are dominated by BB,
611 using BC as a solo tracer to calculate the primary ratio (MAE_p) still works. In scenario B in which BC
612 and primary BrC are independent, using BC alone to determine a single primary MAE_p could lead to
613 a considerable bias in E_{abs} estimation. Alternatively, if a reliable primary BrC tracer is available, the
614 corresponding $MAE_{p,BrC}$ can be determined by MRS. With the knowledge of $MAE_{p,BrC}$ and $MAE_{p,BC}$,
615 light absorption by BC and BrC can be calculated separately and the E_{abs} can be determined using Eq.
616 (13) :

$$617 \quad E_{abs} = \frac{\sigma_{abs,t}}{\sigma_{abs,p,BC} + \sigma_{abs,p,BrC}} = \frac{\sigma_{abs,t}}{MAE_{p,BC} \times EC + MAE_{p,BrC} \times BrC} \quad (13)$$

618 However, the implementation of Eq.13 is challenging due to the complexity in the chemical
619 composition of BrC. For example, a recent study found that the 20 most absorbing BrC chromophores
620 account for ~50% BrC light absorption and there is not a single compound contributing more than 10%
621 (Lin et al., 2016), making it difficult to choose a single compound as the BrC tracer. In addition, time
622 resolved measurement of BrC chromophores has yet to emerge. As a result, for scenario B (sample
623 $AAE > 2$ & primary BrC variations independent of BC), estimation of E_{abs} by MRS is not practical at

624 this stage due to the lack of required input data. Using BC alone to determine a single primary MAE_p
625 could lead to a considerable bias and should be avoided.

626 **6 Conclusions**

627 In this study, a novel statistical approach is proposed and its application on ambient data is
628 demonstrated using one-year hourly OC and EC data coupled with Aethalometer measurements.
629 Unlike conventional E_{abs} determination approaches that require expensive instrumentation (e.g. TD-
630 PAS, VTDMA, SP2), this new approach employs widely deployed instruments (field carbon analyzer
631 and Aethalometer). The key of this new approach involves calculating MAE_p by the Minimum R
632 Squared (MRS) method (Wu and Yu, 2016). The MRS method opens up a new approach to investigate
633 the long-term trend of E_{abs} that was rarely studied by the TD approach. It is found that E_{abs} estimation
634 by MRS is insensitive to systematic biases in EC and σ_{abs} measurements. The annual average
635 $MAE_{p,550}$ estimated by MRS is $13 \text{ m}^2 \text{ g}^{-1}$ and annual average MAE_{550} is $18.75 \pm 6.16 \text{ m}^2 \text{ g}^{-1}$, suggesting
636 an annual average enhancement factor (E_{abs550}) of 1.50 ± 0.48 in the PRD region. This value is within
637 the upper limit of E_{abs} (~ 2) by core-shell Mie simulations considering the typical soot size distribution
638 and coating thickness in the PRD.

639 Both $MAE_{p,550}$ and E_{abs} show distinct seasonal variations, implying the complexity of soot
640 particle mixing state variations in this region. The elevated summertime E_{abs550} in the PRD is found to
641 be associated with the domination of aged air masses from the South China Sea, along with the long-
642 range transport of biomass burning influenced air masses from Southeast Asia. Core-shell size ranges
643 narrowed down by E_{abs550} and $AAE_{470-660}$ constraints suggest that soot particles in the rainy season are
644 likely to have thicker coatings than in the dry season.

645

646 **Data availability**

647 OC, EC, inorganic ions and σ_{abs} data used in this study are available from corresponding authors
648 upon request.

649

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659

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1065 Table 1. Abbreviations.

1066

Abbreviation	Definition
AAE ₄₇₀₋₆₆₀	Absorption Angstrom Exponent between 470 and 660 nm
BB	Biomass burning
BrC	Brown Carbon
D _{core} , D _{shell}	Particle diameter of core/shell
E _{abs550}	Light absorption enhancement factor at 550 nm
σ_{abs550}	Light absorption coefficient at 550 nm
$\sigma_{abs,t}$	Total light absorption coefficient of a coated particle
$\sigma_{abs,p}$	Primary light absorption coefficient attributed to the soot core alone of a coated particle
$\sigma_{abs,c}$	Extra light absorption coefficient due to the lensing effect of coating on the soot core
LII	Laser induced incandescence technique for soot measurement
LWC	Liquid water content
MAE ₅₅₀	Mass absorption efficiency at 550 nm, also known as mass absorption cross-section (MAC)
MAE _{p,550}	Primary MAE of freshly emitted soot particles at 550 nm
MAAP	Multi Angle Absorption Photometer
MOUDI	Micro Orifice Uniform Deposit Impactor
MRS	Minimum R squared method
PAS	Photo acoustic spectrometer
PRD	Pearl River Delta region, China
SP2	Single particle soot photometer
SSA	Single scattering albedo
TD	Thermal denuder
TOA	Thermal optical analysis
TSV	Total spatial variance in backward trajectories cluster analysis

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1068 Table 2. Comparison of MRS application on $(OC/EC)_p$ (for SOC estimation) and MAE_p (for E_{abs} estimation).
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	MRS in EC tracer method for SOC estimation (Wu and Yu, 2016)	MRS in EC tracer method for E_{abs} estimation (this study)
Key parameter of fresh EC particles to be determined	$\left(\frac{OC}{EC}\right)_p = \frac{POC}{EC}$	$MAE_p = \frac{\sigma_{abs,p}}{EC}$
Input quantities for MRS from measurements	OC, EC (tracer)	$\sigma_{abs,t}$, EC (tracer)
Variable to be decoupled by the tracer	$OC = POC + SOC$ $= \left(\frac{OC}{EC}\right)_p \times EC + SOC$	$\sigma_{abs,t} = \sigma_{abs,p} + \sigma_{abs,c}$ $= \frac{\sigma_{abs,p}}{EC} \times EC + \sigma_{abs,c}$
Ambient measurement at its closest to fresh emissions	Minimum R^2 (SOC, EC) $SOC = OC - \left(\frac{OC}{EC}\right)_p \times EC$	Minimum R^2 ($\sigma_{abs,c}$, EC) $\sigma_{abs,c} = \sigma_{abs,t} - MAE_p \times EC$
Graph	<p>Minimum R^2 $(OC/EC)_p = 2.26$</p>	<p>Minimum R^2 $(OC/EC)_p = 13$</p>

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1071 Table 3. Comparison of E_{abs} between various studies.

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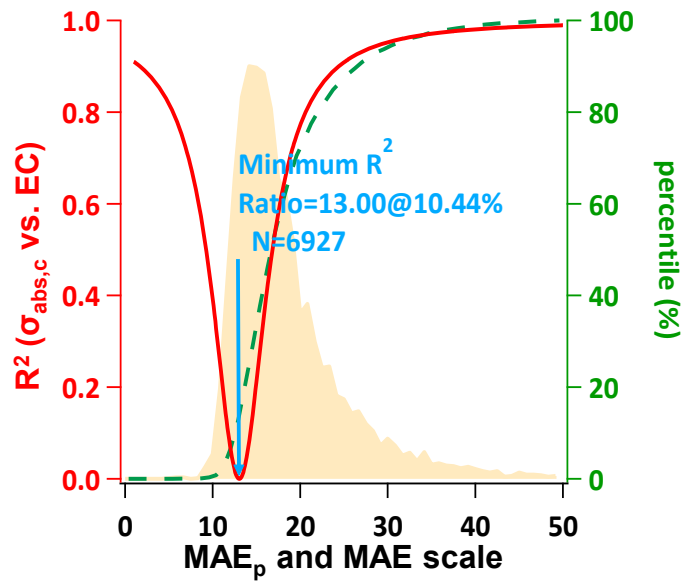
Location	Type	Sampling Duration	λ (nm)	Instrument	E_{abs}	Method	Reference
Guangzhou, China	Suburban	2012.2-2013.1	550	AE+OCEC	1.50±0.48	MAE	This study
Xi'an, China	Urban	2012.12-2013.1	870	PAS	1.8	MAE	(Wang et al., 2014)
Shenzhen, China	Urban	2011.8-9	532	PAS	1.3	MAE	(Lan et al., 2013)
Jinan, China	Urban	2014.2	678	OCEC	2.07 ± 0.72	AFD	(Chen et al., 2017)
Nanjing, China	Suburban	2012.11	532	PAS	1.6	MAE	(Cui et al., 2016a)
Boulder, USA	Forest fire	2010.9	532	PAS	1.38	TD 200°C	(Lack et al., 2012a)
London, UK	Rural	2012.2	781	PAS	1.4	TD 250°C	(Liu et al., 2015)
California, USA	Rural	2010.6	532	PAS	1.06	TD 250°C	(Cappa et al., 2012)
Noto Peninsula, Japan	Rural	2013.4-5	781	PAS	1.22	TD 300°C	(Ueda et al., 2016)
Yuncheng, China	Rural	2014.6-7	678	OCEC	2.25 ± 0.55	AFD	(Cui et al., 2016b)
San Jose, Costa Rica	Rural	2006 winter	1064	SP2	1.3	Mie+SP2	(Schwarz et al., 2008)

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AE: Aethalometer ; OCEC: OCEC analyzer; PAS: photo acoustic spectrometer; SP2: Single particle soot photometer; TD: Thermal denuder AFD: filter filtration-dissolution

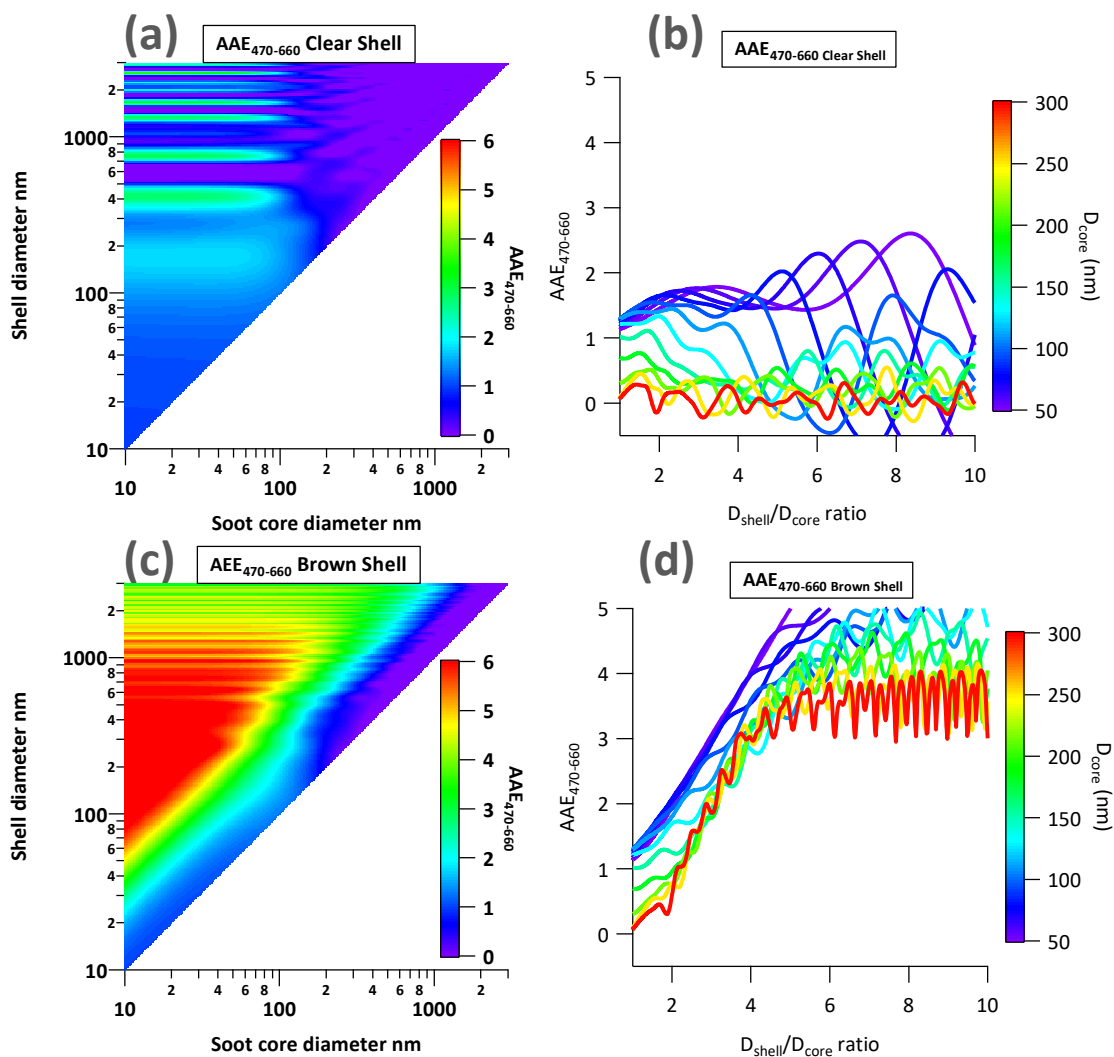
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1077 Figure 1. Minimum R squared (MRS) plot for calculating MAE_p at 550 nm. The red curve is the correlation result between
 1078 $\sigma_{abs,c}$ ($\sigma_{abs,t} - EC * MAE_p$) and EC mass. The shaded area in light tan represents the frequency distribution of observed
 1079 MAE. The dashed green line is the cumulative distribution of observed MAE.



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Figure 2. Mie simulated size dependency of soot particles AAE₄₇₀₋₆₆₀. (a) Combination of different clear shell (y axis) and core diameters (x axis). The color coding represents the AAE₄₇₀₋₆₆₀ of a particle with specific core and clear shell size; (b)

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Cross-sections views of (a). The color coding represents different D_{core} in the range of 50 ~ 300 nm. (c)&(d) Similar to

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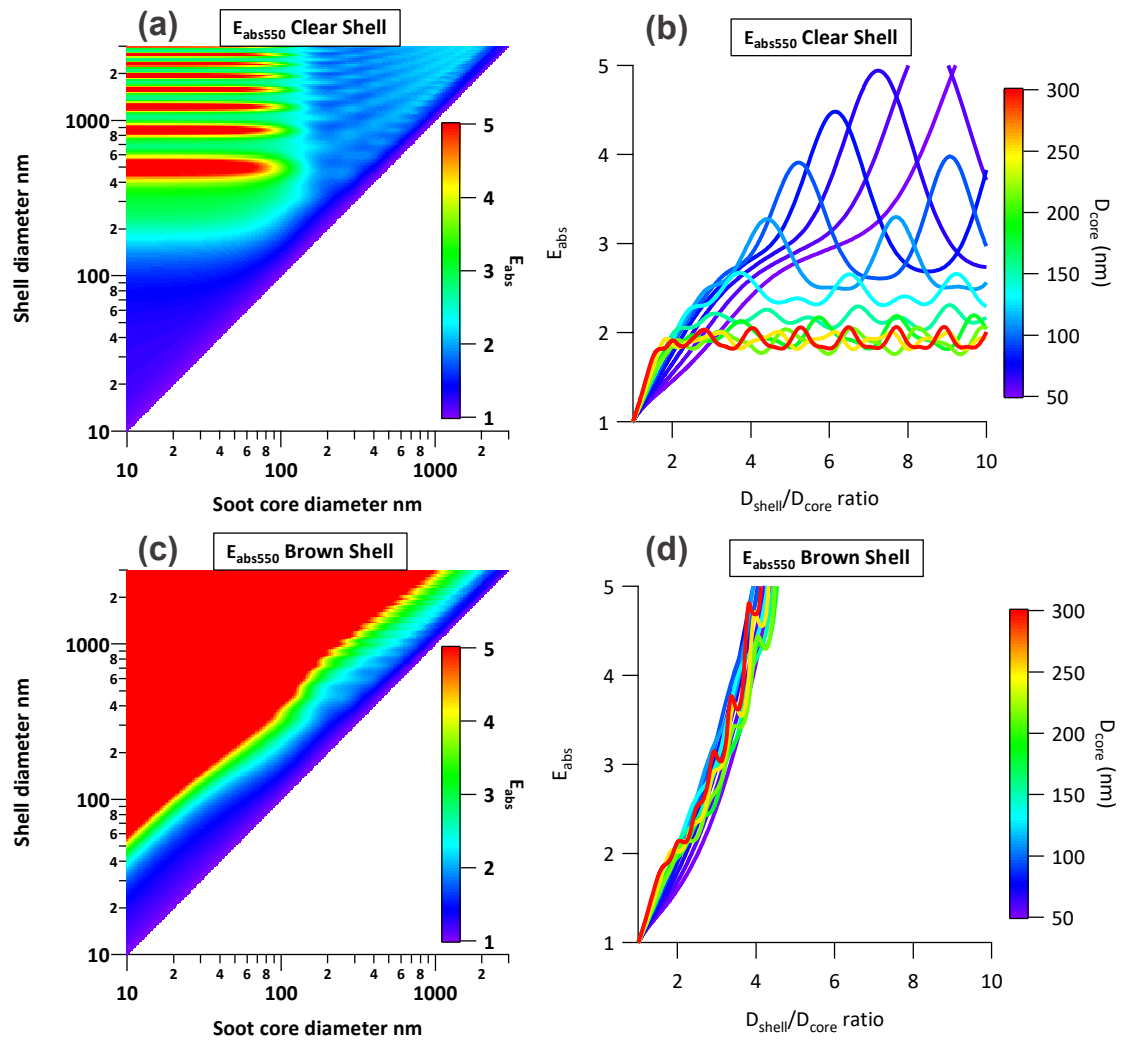
(a)&(b) but from the brown shell scenario.

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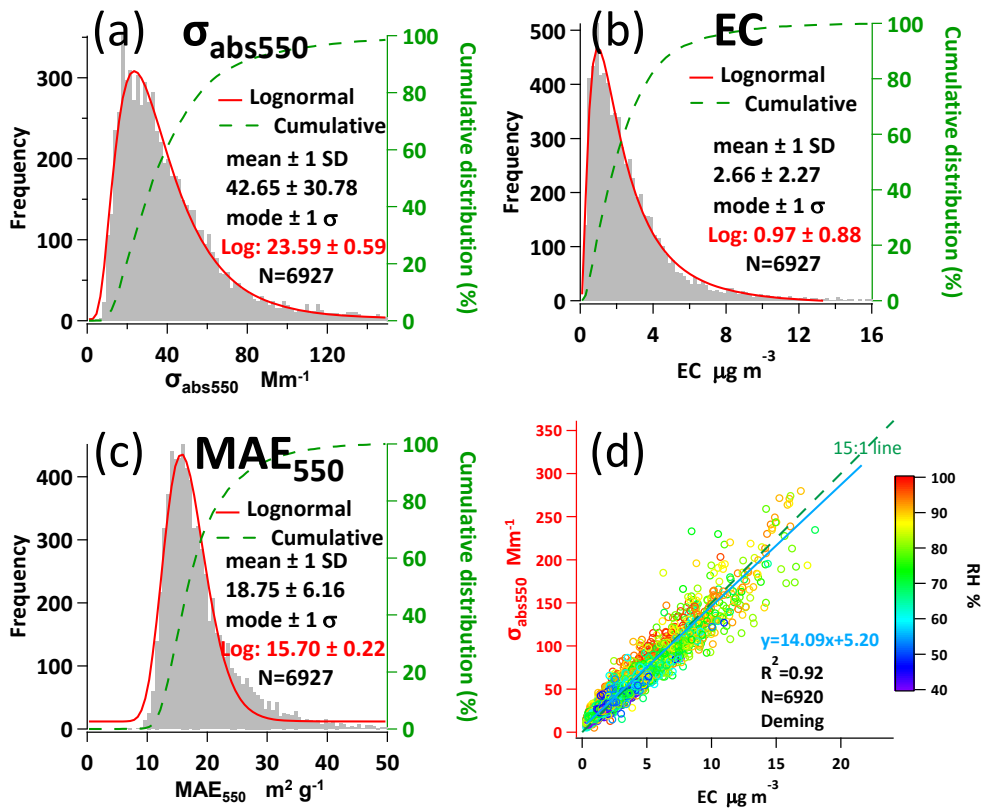
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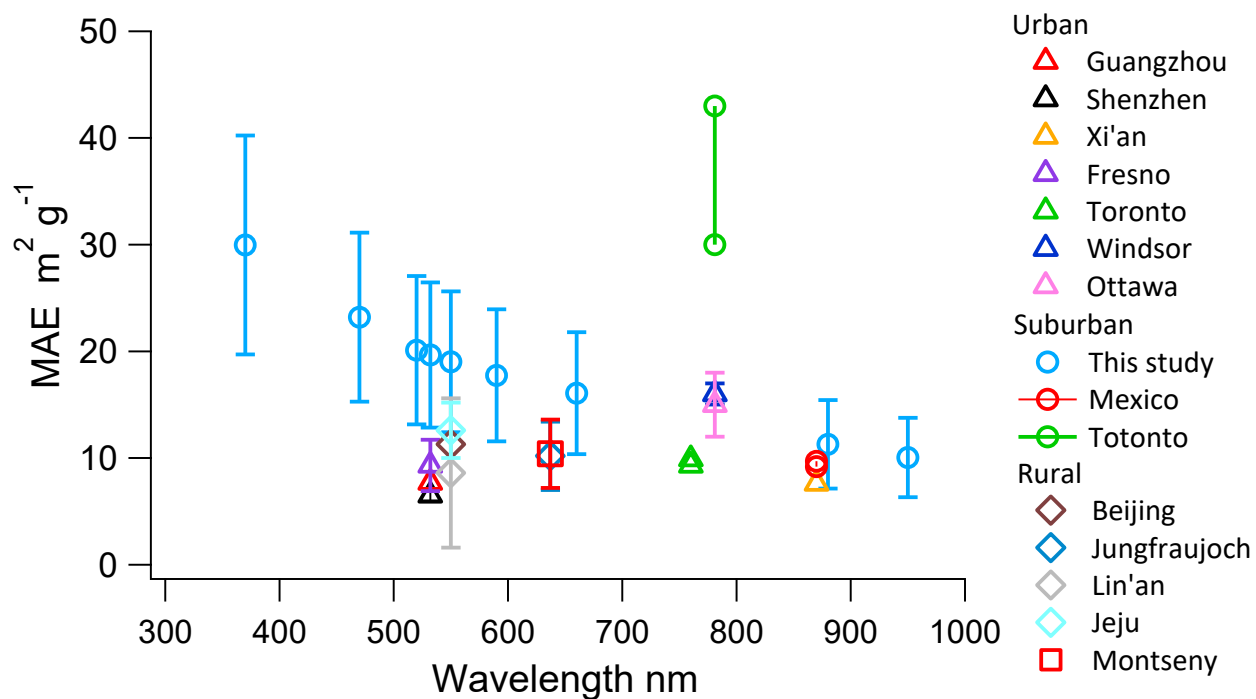
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Figure 3. Mie simulated size dependency of soot particles E_{abs} at wavelength 550 nm. (a) Combination of different clear shell (y axis) and core diameters (x axis). The color coding represents the E_{abs} of a particle with specific core and clear shell size; (b) Cross-sections views of (a). The color coding represents different D_{core} in the range of 50 – 300 nm. (c)&(d) Similar to (a)&(b) but from the brown shell scenario.



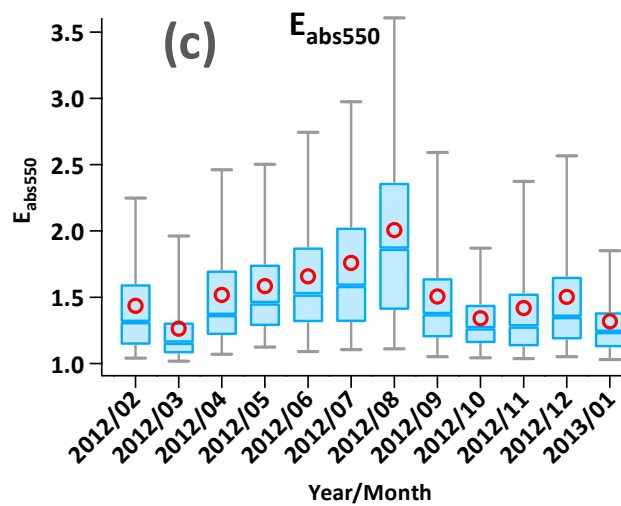
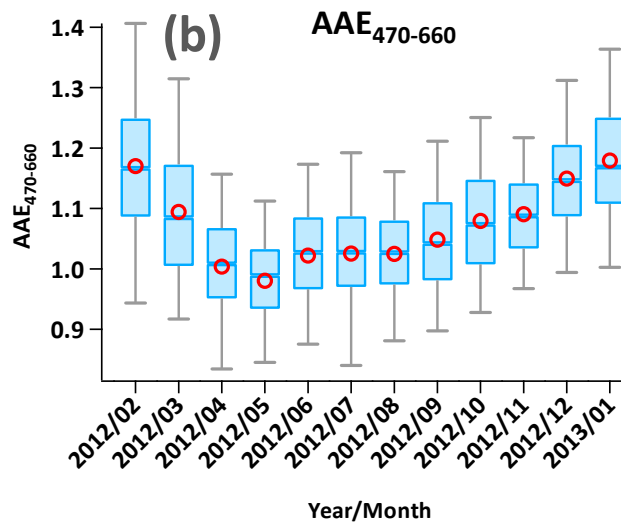
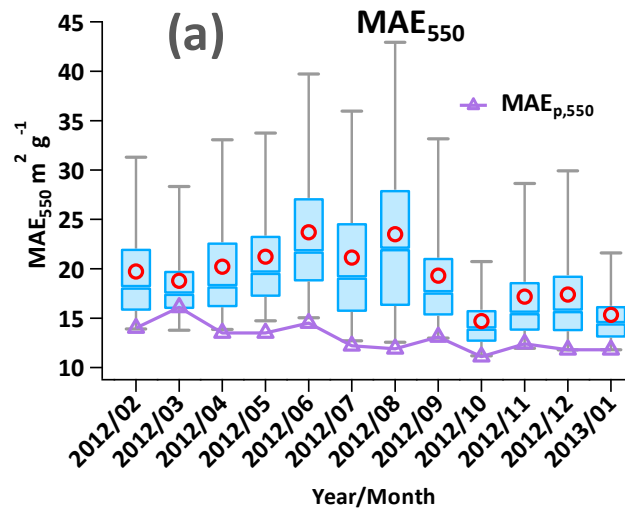
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1097 Figure 4. Measured annual statistics of σ_{abs550} , EC and MAE₅₅₀. (a) Annual frequency distribution of light absorption at
 1098 550 nm. The red curve represents the fitting line for a log-normal distribution. (b) Annual frequency distribution of EC
 1099 mass concentration (c) Frequency distribution of Mass absorption efficiency (MAE) at 550 nm. (d) Scatter plot of light
 1100 absorption (550 nm) and EC mass. The slope represents MAE₅₅₀. The blue regression line is by Deming regression. The
 1101 color coding represents RH.
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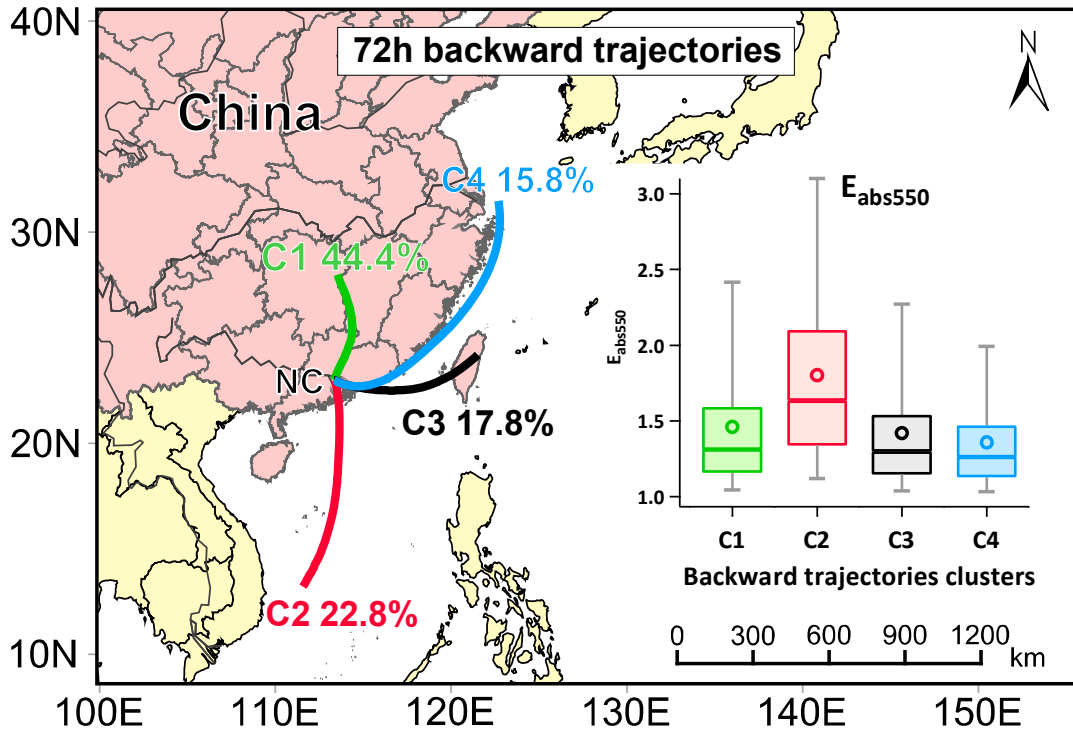
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1104 Figure 5. Comparison of spectral MAE measurements from this study with previous studies. Triangle, circle and rhombus
 1105 represent urban, suburban and rural respectively. Details and reference can be found in Table S1. The whiskers represent
 1106 one standard deviation.



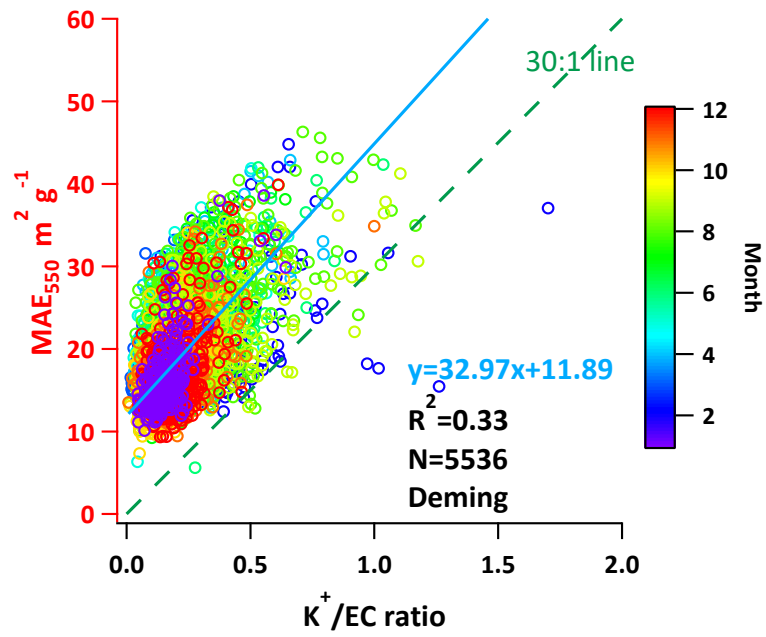
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1108 Figure 6. Measured monthly variations of (a) MAE_{550} , the purple line represents $MAE_{p,550}$ estimated by MRS (b) AAE_{470-}
 1109 660 and (c) E_{abs550} . Red circles represent the monthly average. The line inside the box indicates the monthly median. Upper
 1110 and lower boundaries of the box represent the 75th and the 25th percentiles; the whiskers above and below each box represent
 1111 the 95th and 5th percentiles.



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1113 Figure 7. Average backward trajectories arriving at 100 m at NC site for four clusters (2012 Feb - 2013 Jan). E_{abs550} by
 1114 different clusters are shown in the box plot.



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1116 Figure 8. MAE₅₅₀ dependency on biomass burning indicator K^+/EC ratio. The color coding represents months. The intercept
 1117 represents MAE without biomass burning effect. The 30:1 line serves as a reference line with an integer slope that
 1118 is close to the regressed slope through the origin.