

## Point-by-point response to review comments on manuscript acp-2017-582 “Quantifying black carbon light absorption enhancement by a novel statistical approach”

By Cheng Wu et al.

We thank the two anonymous reviewers for their constructive comments to improve the manuscript. Our point-by-point responses to the review comments are listed below. Changes to the manuscript are marked in blue in the revised manuscript. The marked manuscript is submitted together with this response document.

### Anonymous Referee #2

**R2-Q1.** The manuscript presents a statistical analysis of black carbon light absorption enhancement based on observations made over roughly one year using a filter-based absorption instrument and a thermal-optical analysis OC/EC analyzer. They derive the absorption enhancement ( $E_{\text{abs}}$ ) from the total mass absorption efficiency measured from the ratio of the absorption and EC measurements to that estimated for bare BC particles. To determine the bare BC MAE the authors employ a method that searches for an MAE value based on an assumed independence between EC and absorption due to BC coatings. While this approach presents a potential alternative to more expensive and labor-intensive methods to examine the important topic of BC light absorption enhancement, it is not clear to me that it works based on the information presented in the manuscript. First, while it does compare  $E_{\text{abs}}$  derived using this method to previously reported values using other techniques, there is not a direct comparison between the results obtained using this technique and more established methods for the same site. As pointed out by Liu et al. (2015), the “influence of coatings on BC absorption may be source and regionally specific”, so it is difficult to draw too much confidence in the approach based on similarities with other locations. Second, it is not obvious that the measurement of EC is independent of the amount of light absorbed by BC coatings. While it is true that the mass of primary BC should be independent of coating absorption, EC is an operationally defined quantity, and depending on how coatings interact with its measurement during thermal optical analysis, could in fact have a relationship with coatings and light absorption due to coatings. For these two principle reasons I do not recommend the manuscripts publication in ACP in its current form.

### Author’s Response:

- (1) A comparison would definitely be helpful for the verification of the MRS approach in the future. However, as the reviewer has pointed out, TD approach may not be a perfect “time machine” to reverse the aging process for  $E_{\text{abs}}$  determination. The definition of  $\text{MAE}_p$  by MRS is different from the  $\text{MAE}_p$  by TD.  $\text{MAE}_p$  by MRS reflect the MAE of soot particles at the emission source, while  $\text{MAE}_p$  by TD reflect the MAE of bare soot particles. First, the morphology of thermally denuded soot particles (collapsed globules) is different from the morphology of freshly emitted BC particles (fractal-like aggregates). Second, most of the coatings are removed for TD denuded soot particles, but freshly emitted soot particles usually come with a thin coating due to the temperature drop from engine to the ambient air. As a result, the  $\text{MAE}_p$  by TD approach is expected to be lower than the  $\text{MAE}_p$  of emission source retrieved by MRS. In this sense, it may be difficult to extract quantitative information through such comparisons due to the difference in  $\text{MAE}_p$  definition.
- (2) We agree with the reviewer that coating can affect EC determination by the thermal-optical analysis. Lee et al. (2007) used artificially fabricated EC samples with OC coatings to evaluate the performance of OC/EC analysis. Biases were observed but the results are linearly correlated ( $R^2 > 0.9$ ) with the true OC and EC values. As shown in Figure R-5 and discussions in the following response to R2-Q5, the  $E_{\text{abs}}$  estimation by MRS is insensitive to systematic biases. The MRS approach mainly rely on correlation analysis, the systematic biases are not a big concern as long as the measured data correlated well with the true values.

Please see below for point-by-point response to reviewers’ comments.

## General Comments

### **R2-Q2.** RH impacts

The manuscript does not clearly state whether air was dried prior to sampling with the instruments. Section 4.5, which discusses impacts of RH on the observations, implies that it was not. The methods section states that 2.5  $\mu\text{m}$  cyclones were used upstream of the Aethalometer and Sunset instruments, but does not specify a size cut for the nephelometer or MARGA instruments. If the nephelometer and MARGA instruments did not have a size cut, it is extremely difficult to compare results from those instruments to those from the AE-31 and Sunset due to potential differences from coarse mode particle contributions.

### **Author's Response:**

Both nephelometer and MARGA are equipped with a  $\text{PM}_{2.5}$  inlet to remove the coarse particles. The following text is added in section 2 to improve the clarity.

Both instruments are equipped with a  $\text{PM}_{2.5}$  inlet to remove the coarse particles.

**R2-Q3.** If the air was not dried prior to sampling more details need to be provided regarding the effective RH for the optical instruments (AE-31 and nephelometer), which may be different from ambient RH due to temperature differences between the ambient air and the instruments. In addition, if air was sampled at near ambient RH it could have a number of complicating factors on the subsequent analysis. Filter-based absorption instruments, as the authors acknowledge, are affected by artifacts, including the scattering and shadowing of light in the filter matrix by filter fibers and particles embedded on the filter. Varying sample RH can change the filter artifacts in difficult to account for ways. For example, if hygroscopic particles are present on the filter they can swell or shrink depending on the sample RH, changing the light transmission properties of the filter, artifacts affecting the absorption measurement and possibly the absorption measurement itself depending on how RH interacts with the correction methods applied. The manuscript should include a discussion of how well the correction methods applied can account for changes arising from changes in RH. In addition, the effective size cut of the cyclone will be affected by RH in that the aerosol sampled will be different depending on how much water is associated with it. This effect may be small depending on the aerosol distribution and makeup of the light-absorbing particles, but should be addressed.

**Author's Response:** Thanks for the very insightful comments. We fully agree that RH could be a source of uncertainty for filter based  $\sigma_{abs}$  measurement. Arnott et al. (2003) found particulate soot absorption photometer (PSAP) showed erratic response as RH change. Schmid et al. (2006) observed dependency of PSAP  $\sigma_{abs}$  on RH, but the effect of RH on Aethalometer performance is negligible. Inter-comparison studies shown that with proper corrections, Aethalometer  $\sigma_{abs}$  agrees well with PAS (Ajtai et al., 2011). During the inter-comparison study of Aethalometer (AE-16) and PAS in Guangzhou (Wu et al., 2009), a good correlation was found ( $R^2=0.96$ ) as shown in Figure R1 or Figure S1. These filed comparison results imply that Aethalometer results are linearly correlated with PAS and RH has limited effect on Aethalometer measurements. Please refer to our response to R1-Q2 for the revisions made in the manuscript.

**R2-Q4.** Another potential issue related to RH affects some of the results presented in Section 4.5. First, it is not clear if the RH reported is ambient RH or instrument RH. If ambient RH, it needs to be established that the RH at the location of the measurement (the filter) is also at the same RH as the ambient. In addition, it is not clear that the absorption measured on the AE-31 filter represents the absorption in the ambient air. For example, if BC becomes coated on the filter (which can happen in the presence of liquid organic aerosol (e.g. Subramanian et al., 2007) it is not clear how water uptake by those organic films might alter the absorption by particles that were not originally coated. Non-BC containing material can also take up water, alter the optical properties of the filter, and change the apparent absorption attributed to BC. Also, since the AE absorption measurement is based on the change in attenuation over time if the hygroscopic properties of particles change with time the apparent

absorption could be affected. For example, consider a situation where BC with a hygroscopic coating has been sampled onto the filter for some period of time. As the sample RH begins to decrease the water will evaporate from the filter, likely leading to a decrease in attenuation (there is a reduced enhancement of absorption, for example). At the same time consider what happens if additional BC is sampled to the filter. This acts to increase attenuation, opposing the RH effect. As a result the BC measured during this time would be underestimated (its effect on attenuation is countered to some degree by the opposing RH effect). Note that the magnitude of these impacts would be difficult to account for, in that they depend on the hygroscopicity of the material on the filter and how that material and associated water interacts in the filter matrix.

**Author's Response:** The RH used in this study is ambient RH. We agree with the reviewer that RH at the filter of Aethalometer is different from the ambient RH. The RH in the optical chamber of Aethalometer may be lower than the ambient RH due to the slightly elevated temperature inside the instrument. The magnitude of RH difference is similar between different instruments: 20% for the Aethalometer (Schmid et al., 2006) and 15% for the nephelometer (Guyon et al., 2004). Although the RH in the Aethalometer optical chamber was not measured in this study, its level was expected to be slightly lower than the ambient RH. Schmid et al. (2006) found PSAP  $\sigma_{abs}$  showed dependency on RH, but the effect of RH on Aethalometer performance was negligible. Cappa et al. (2008) found  $\sigma_{abs}$  by PSAP maintained high linearity with PAS even under high RH conditions (65-91%). Inter-comparison studies shown that with appropriated corrections, Aethalometer  $\sigma_{abs}$  agrees well with PAS (Ajtai et al., 2011). During the inter-comparison study of Aethalometer (AE-16) and PAS in Guangzhou (Wu et al., 2009), good correlation was found ( $R^2=0.96$ ) as shown in Figure S1. These comparison results imply that Aethalometer results are linearly correlated with PAS and RH has limited interference on Aethalometer measurements.

**R2-Q5.** Measurement biases and impacts on the analysis Biases in filter-based absorption and OC/EC measurements are mostly associated with other materials mixed and/or co-sampled with BC/EC, so it is likely that errors in the measurements have a systematic relationship. For example, the presence of organic aerosol and organic films has been linked to both biases in filter-based measurements (e.g., Lack et al., 2008) as well as potential impacts on EC measurement via pyrolyzed carbon correction (Subramanian et al., 2007). The manuscript needs a detailed discussion of how correlations in biases between the methods affect the retrieved MAE<sub>p</sub> values and resulting E<sub>abs</sub> calculations. Since the manuscript is seeking to establish a new method this potential issue needs much more attention. Several specific comments below are related to this general concern.

**Author's Response:** The bead-like particles on filter fibers could interfere both Aethalometer and OC/EC measurements. Since the study by Subramanian et al. (2007) was based on source samples from low temperature BB emissions, how this bead-like particles could affect ambient measurements depends on their fractional contribution to PM<sub>2.5</sub>. Cappa et al. (2008) found that the light absorption enhancement factor for the PSAP relative to the PAS due to the presence of externally mixed organic aerosols is proportional to the OA-to-soot ratio, which means this bias is systematic. We conduct two tests to investigate the effect of systematic EC and  $\sigma_{abs}$  bias on E<sub>abs</sub> estimation. As shown in Figure R-5, the MRS approach is insensitive to systematic bias in EC and  $\sigma_{abs}$  measurements. Details of the tests are added in section 4.1 and also shown below.

To investigate the performance of the MRS approach in response to systematic bias in EC and  $\sigma_{abs}$ , two simple tests are conducted as shown in Figures S9 and S10 by adding systematic biases to  $\sigma_{abs550}$  and EC. Test A represents a situation when  $\sigma_{abs}$  is overestimated and EC is underestimated. The biased data are marked as  $\sigma'_{abs550}$  and EC' respectively, as shown below:

$$\sigma'_{abs550} = \sigma_{abs550} \times 2 \quad (8)$$

$$EC' = EC \times 0.7 \quad (9)$$

As a result, the average  $MAE_{550}$  changed from 18.75 to 53.58  $m^2 g^{-1}$  and  $MAE_p$  changed from 13 to 37  $m^2 g^{-1}$  (Figure S9). However,  $E_{abs}$  by ratio of averages remain the same (1.44).

In Test B, EC by different TOA protocols are compared to investigate the effect of different EC determination approaches while  $\sigma_{abs550}$  remains unchanged. EC by IMPROVE TOR protocol is calculated from NIOSH TOT EC following an empirical formula for suburban sites derived from a 3-year OCEC dataset in PRD (Wu et al., 2016):

$$EC_{IMP\_TOR} = 2.63 \times EC_{NSH\_TOT} + 0.05 \quad (10)$$

As shown in Figure S10,  $MAE_{550}$  changed from 18.75 to 7.02  $m^2 g^{-1}$  and  $MAE_p$  changed from 13 to 5  $m^2 g^{-1}$ , but  $E_{abs}$  remain almost the same (1.40). Result of Test B implies that although EC is operationally defined, the discrepancy of EC between TOA protocols did not weaken the role of EC serving as a tracer for primary emissions in MRS application. These examples demonstrate that systematic biases in  $\sigma_{abs550}$  and EC have no effects on  $E_{abs}$  estimation by the MRS approach.

# Test A

Original data

Systematically biased data

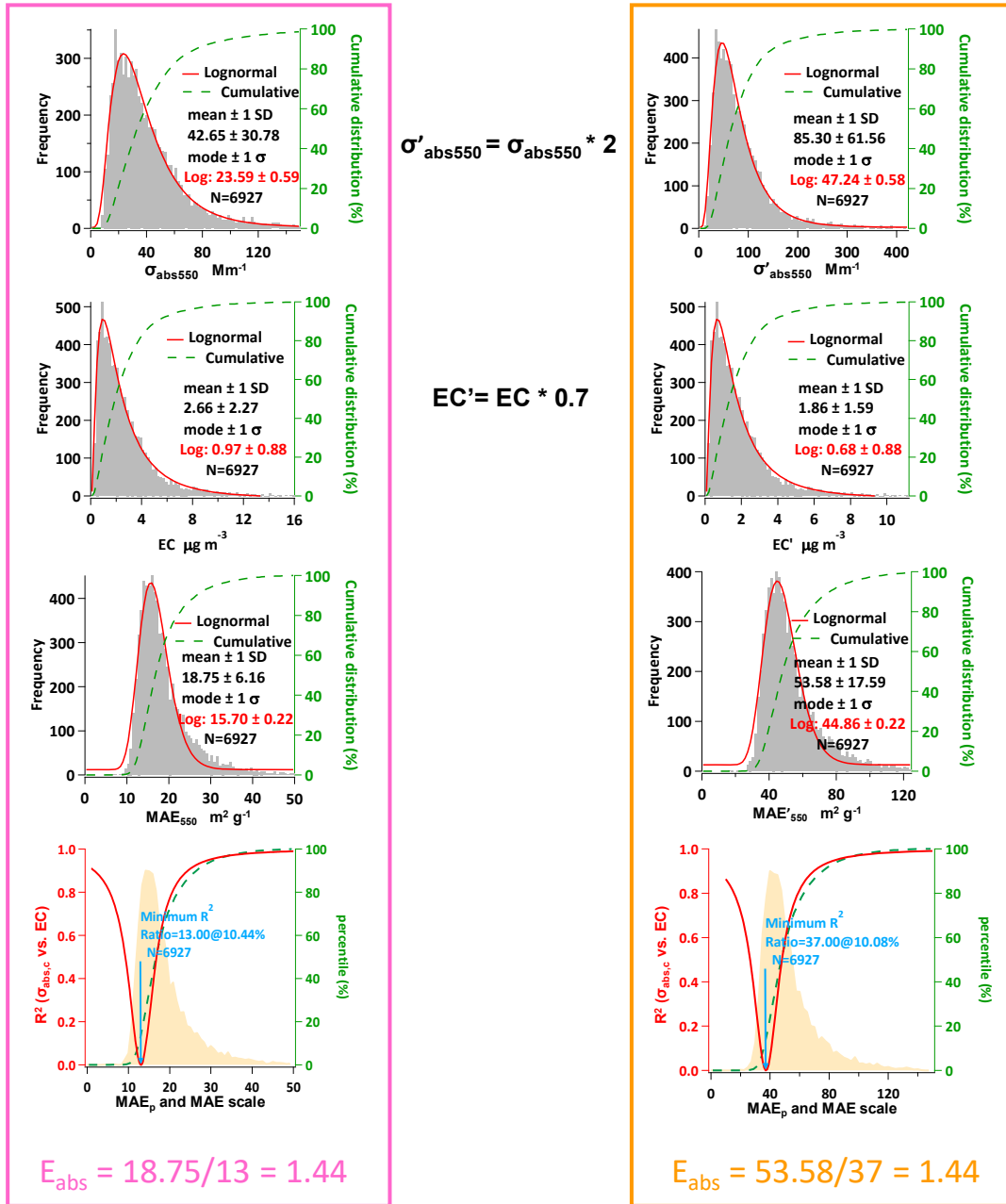


Figure R-5. Comparison of  $E_{\text{abs}}$  from original data and systematically biased data (Test A).

## Specific Comments

**R2-Q6.** 78-80: Missing from the list of field studies is Cappa et al. (2012), which found a weak enhancement for observations in California, US.

**Author's Response:** Suggestion taken. Cappa et al. (2012) was added.

**R2-Q7.** 85: Suggest changing to “The TD approach is briefly discussed here.”

**Author's Response:** Revision made.

**R2-Q8.** 90: Some good points are raised in this discussion, but a few things are lacking. First, another major reason for use of PAS systems in the approach is that the technique does not have artifacts associated with filter-based methods for measuring light absorption, thus it provides an unambiguous measure of the light absorption coefficient in both heated and un-heated states. Also missing from the discussion is potential differences in BC core morphology being different for “fresh” versus “aged with coating removed” conditions, or in other words, the TD may not be a perfect “time machine” to reverse the aging process and determine the effects of BC aging.

**Author's Response:** Thanks for the suggestions. The following content is added in the main text.

As an in-situ technique, PAS eliminates the artifacts associated with filter-based methods (Weingartner et al., 2003;Coen et al., 2010) and is often considered as the reference instrument for light absorption coefficient determination (Arnott et al., 2003;Arnott et al., 2005).

It's also worth noting that  $MAE_p$  by the TD approach is different from the  $MAE_p$  at the emission source. First, the morphology of thermally denuded BC particles (compact aggregates) is different from that of freshly emitted BC particles (chain-like aggregates). Second, most of the coatings is removed with the TD denuded soot particles, but freshly emitted soot particles usually come with a thin coating of OC formed from condensation of OC vapors as the temperature drops from engine to the ambient air. As a result, the  $MAE_p$  by TD approach is expected to be lower than the  $MAE_p$  of emission source. In this sense, the TD approach may not be a perfect “time machine” to reverse the aging process for  $E_{abs}$  determination.

**R2-Q9.** 100-101: I realize this is not the focus of the manuscript, but it is worth noting here the logic presented here assumes a linear relationship between MAE and coated soot particle number fraction measured by SP2. This may not be the case due to limitations in the size of BC particles measured by the SP2, its ability to measure thin coatings, the potential impact of thin coatings on MA, the assumption of core-shell morphology, and the nature of the relationship between the fraction of coated particles (a parameter which ignores coating thickness, the physical driver of the absorption enhancement) and MAE.

**Author's Response:** The following text is added in the main text for clarification.

It is worth noting that this approach provides only a rough approximation of  $E_{abs}$  since the parameter used here (coated soot particles number fraction) ignores other main drivers of light absorption enhancement (e.g. coating thickness). As a result, this approach is only valid for a period of measurements, for which coating thickness is

relatively constant and the MAE variations are dominated by coated soot particles number fraction.

**R2-Q10.** 133: What does the 5% uncertainty refer to? Is this before or after corrections? Are the uncertainties in the corrections as low as 5%? Also, the manuscript gives an uncertainty of 24% for the ECOC analyzer here, but then later states (line 153) up to a factor of 5 differences in EC measured by different protocols are possible. How is this to be reconciled?

**Author's Response:** The 5% measurement uncertainty refers to instrument precision. An uncertainty of 24% for the ECOC analyzer is the instrument precision when NIOSH protocol is applied for TOA. The 5 times differences in EC measured by different protocols arises from inter-protocol discrepancy. It's a concept different from the instrument precision. To avoid confusion, we change the term "measurement uncertainty" to "measurement precision" throughout the manuscript.

**R2-Q11.** 166-169: The argument given here is only true if the errors in absorption measurement and EC mass measurement are independent and random. If some factor causes a positive bias in the filter-based absorption but a negative bias in the EC measurement they will not cancel, but instead there will be an apparent and potentially false apparent absorption enhancement. Or consider a situation where some co-sampled material affects the OC/EC split but not the filter-based absorption measurement, leading to an apparent change in  $E_{abs}$ .

**Author's Response:** We conduct two tests to investigate the effect of systematic EC and  $\sigma_{abs}$  bias on  $E_{abs}$  estimation. As shown in Figure R-5 and discussions in the response to R2-Q5, the MRS approach is insensitive to systematic bias in EC and  $\sigma_{abs}$  measurements.

Following contents are added in the manuscript.

### **In section 2.1**

Systematic bias in MAE (e.g. overestimation of  $\sigma_{abs}$  and variability of EC mass by different TOA protocols) discussed above have little effect on  $E_{abs}$  estimation by MRS. As shown in Eq. 3,  $E_{abs}$  is the ratio of  $MAE_t$  to  $MAE_p$  or  $\sigma_{abs,t}$  to  $\sigma_{abs,p}$ , thus most of the bias in EC mass or  $\sigma_{abs}$  is cancelled out during the  $E_{abs}$  calculation. More details are discussed in section 4.1.

Please refer to our response to R2-Q5 for contents added in section 4.1.

**R2-Q12.** 187-188: Since EC is an operationally defined parameter based on a measurement technique there is not necessarily an inherent interdependency between it and the absorption due to coatings gained following emission.

**Author's Response:** It is true that EC is operationally defined. But studies have shown that EC by different protocols correlate very well (Chow et al., 2001; Chow et al., 2004; Wu et al., 2012; Wu et al., 2016). The discrepancy of EC between thermal-optical analysis protocols did not weaken the role of EC to serve as a tracer for primary emission in MRS application, as shown in the Test B in the response to R2-Q5. On the other hand, extra absorption due to coating is associate with secondary process after emission. The variations of primary emission are relatively independent to the variations of secondary process.

**R2-Q13.** Section 3.2: The results from the Mie simulations are not a new contribution and main points drawn from the discussion could be drawn from previous studies given in the literature (e.g., Lack and Cappa, 2010; Bond et al., 2006).

**Author's Response:** The main purpose of including Mie simulations in this study is to help readers to understand the viability of  $E_{\text{abs}}$  and AAE from a theoretical perspective and their dependence on different core/shell diameter combinations. Citing main points from literature is one way but that could be distracting. From a reader point of view, digging the detail information from the literature might not be as convenient as reading the Mie simulations in this study. In addition, Mie simulations figures shown here are specifically plotted to support the later discussion of the measurement results. We feel that the inclusion of Mie simulations provides a smooth one-stop reading experience.

**R2-Q14.** 260: Should note here that the AAE observed in measurement studies also includes absorption by potentially externally mixed brown carbon particles, not just those present in the form of shells on BC particles.

**Author's Response:** Suggestion taken. The corresponding content has been revised as follows.

These high AAE results are consistent with the previous model study (Lack and Cappa, 2010) and could partially explain the high AAE observed in measurement studies (Kirchstetter et al., 2004; Hoffer et al., 2006), since the presence of externally mixed BrC particles also contribute to the wavelength dependent light absorption.

**R2-Q15.** 295: Suggest changing “concentrations” to “modes”.

**Author's Response:** Revision applied.

**R2-Q16.** 299-302: Minor point here, but the MOUDI data also reflect differences in BC core equivalent diameters measured using LII and aerodynamic diameter of BC, so we would not necessarily expect similar sizing, even for uncoated BC. Worth mentioning here, though I do not disagree with the main argument given here.

**Author's Response:** Thanks for the clarification. The following content is added to remind readers the difference when comparing the sizing between LII and MOUDI.

BC sizing by LII is based on volume equivalent diameter (VED), while MOUDI is based on aerodynamic diameter. As a result, these two techniques do not necessarily yield similar sizes, even for the bare soot particles. The conversion between these two types of diameters involves the knowledge of particle density and morphology (drag force).

**R2-Q17.** 303: Please state whether Tan et al. refer to BC only diameter or mixed particle diameter here.

**Author's Response:** Tan et al. refers to coated BC diameter. The corresponding content has been revised as follows to improve the clarity.

A recent closure study on BC mixing state in the PRD region suggests  $\sigma_{\text{abs}}$  is dominated by coated soot particles in the range of 300~400 nm (Tan et al., 2016).

**R2-Q18.** 314: I am curious as to why the 470 and 660 nm wavelength pairs were used to quantify AAE. Brown carbon tends to show much stronger impact at shorter wavelengths. Could the authors please include the AAE determined for the UV and 880 nm channels also measured by the AE-31?



**Author's Response:** The 470 and 660 nm wavelength pairs were used to represent AAE in the visible range.  $AAE_{370-880}$  ( $1.13 \pm 0.13$ ) is added as shown in Figure S8b, which is slightly higher than  $AAE_{470-660}$  ( $1.09 \pm 0.13$ ) shown in Figure S8a.

**R2-Q19.** 321: I assume  $MAE_p$  should be  $MAE_{p,550}$  based on Table 1?

**Author's Response:** Thanks for pointing out. Revision made.

**R2-Q20.** 325-328: The Cappa et al. observation showing weak enhancement should also be included in this summary.

**Author's Response:** Suggestion taken. Cappa et al. (2012) was added.

**R2-Q21.** 355: Use of “significantly” implies a statistically significant difference between the clusters. Please provide uncertainty estimates and confidence levels if this is intended, or omit.

**Author's Response:** Results of Wilcoxon-Mann-Whitney tests are included in SI to support the statement and also shown below. Wilcoxon-Mann-Whitney tests between C1&C2, C2&C3 and C2&C4 all show  $P < 0.01$ , indicating that the mean of C2 is significantly higher than C1, C3 and C4.

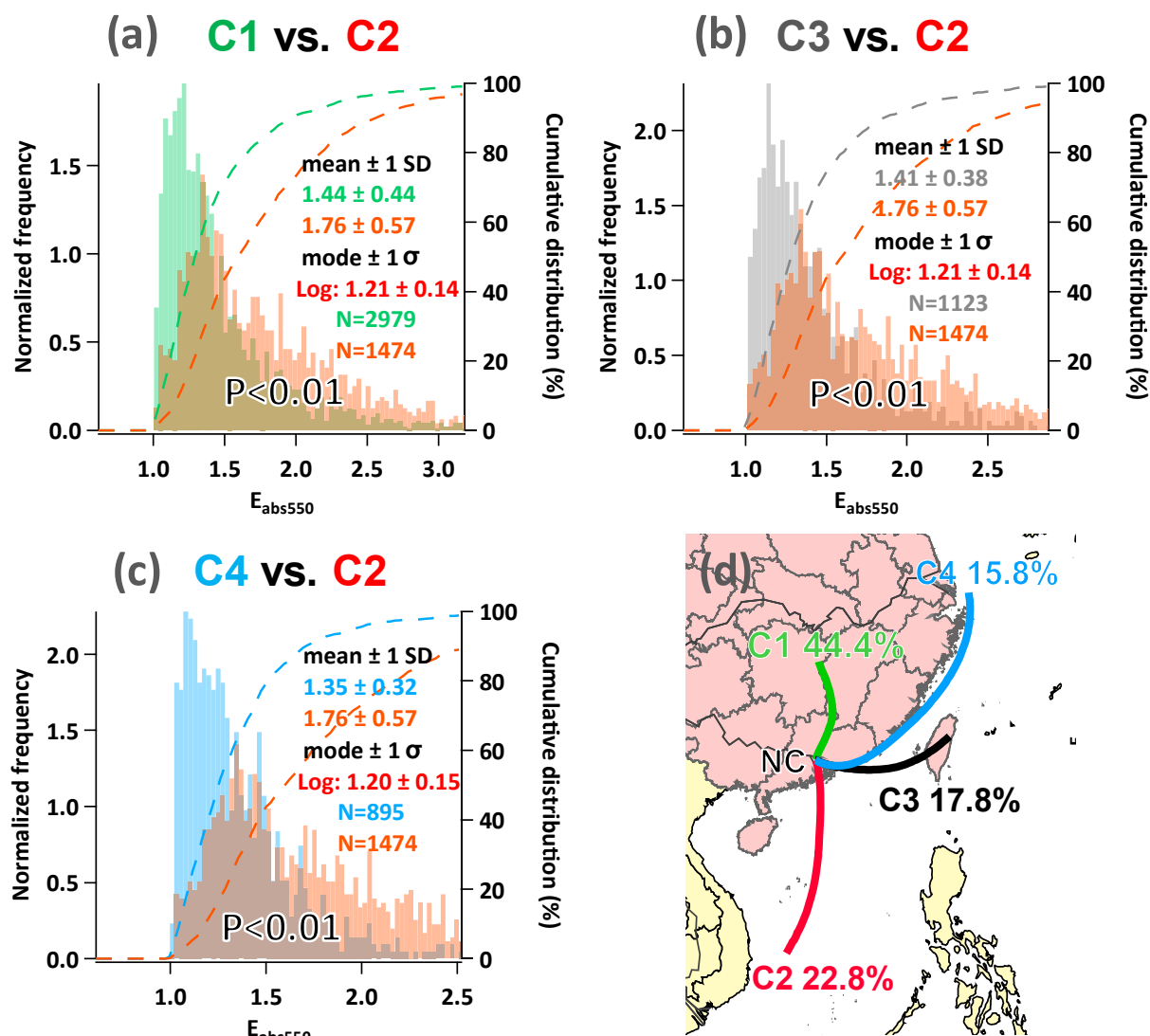


Figure R-6 Frequency distributions of  $E_{abs550}$  by different air mass clusters.

**R2-Q22.** 360-361: I am curious how much the air mass trajectories were influenced by precipitation during the monsoon period/rainy season. It seems like aged/coated BC likely having higher  $E_{abs}$  would be more susceptible to removal compared to less-aged/coated BC with lower  $E_{abs}$ ?

**Author's Response:** It's difficult to directly evaluate the rain effect on  $E_{abs550}$  for corresponding air mass trajectories since the measurement is only conducted at the end point of air mass trajectories. Alternatively, we compare the  $E_{abs550}$  before and during rain for 49 rain events from the yearlong measurements. Precipitation of 49 rain events as well as the monthly distributions are shown in Figure R7a. As for subtropical region, precipitation events are dominated in spring and summer time. EC concentration is only 43% during rain comparing before rain. However, as shown in Figure R7c,  $E_{abs550}$  values are similar before and during rain as indicated by the unity slope. These results imply that  $E_{abs550}$  is not very sensitive to rain effect in this study.

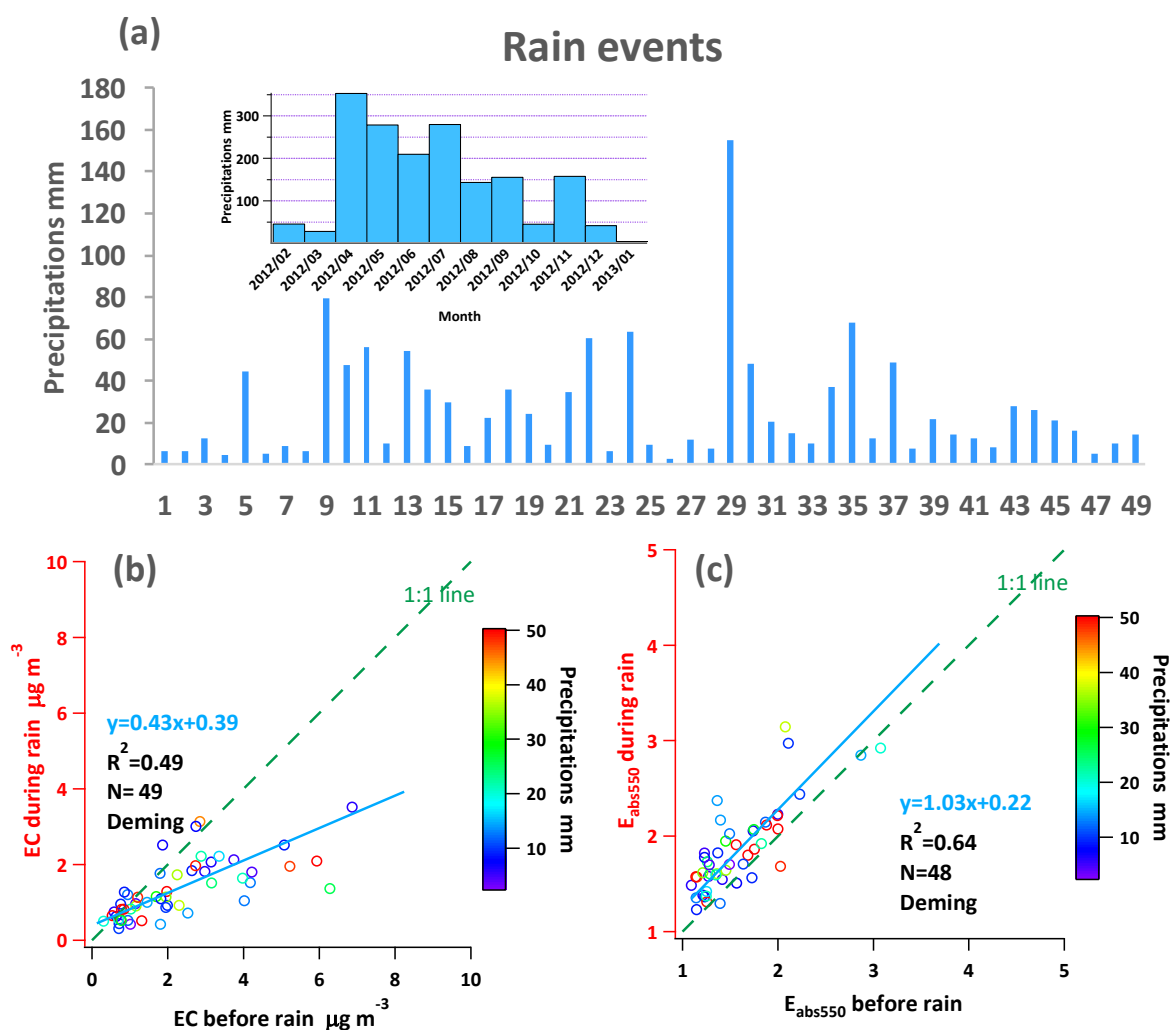


Figure R-7 Effect of precipitations. (a) Precipitations of 49 rain events and monthly distributions. (b) Scatter plot of EC before and during rain events. (c) Scatter plot of  $E_{abs550}$  before and during rain events.

**R2-Q23.** 375-376: Could the authors clarify the reasoning given here? First, I believe the authors mean the  $MAE_p$  determined in section 4.1, not 4.3, correct? I think this is arguing that because  $MAE_p$  is, on average, higher than the MAE observed in the absence of BB influence there must be a large amount of BB influence on BC at this location? If so, I think some caution or caveats should be included in the discussion, since the BB influence tracer is not independent of the EC measurement. Further, I would

expect there to be a much stronger effect on AAE if BB was such an important BC source, yet the later sections establish that AAE does not show a response.

**Author's Response:** Thanks for suggestion. The corresponding content has been revised as follows:

During the rainy season when oceanic wind prevails, BC from BB emission in Southeast Asia can reach PRD through long range transport (LRT), resulting in an elevated  $K^+/EC$  ratio and  $MAE_{550}$ . The Deming regression intercept (11.89) in Figure 8 represents the MAE without the BB effect. This non-BB  $MAE_{550}$  ( $11.89 \text{ m}^2 \text{ g}^{-1}$ ) is only slightly lower than  $MAE_{p,550}$  ( $13 \text{ m}^2 \text{ g}^{-1}$ ) obtained in section 4.3, implying that a large fraction of  $MAE_{p,550}$  could not be explained by the BB source. Additional evidence was obtained through examining regression relationships of  $MAE_{p,550}$  with  $K^+/EC$  month-by-month (Figure S17b). Correlation of monthly  $MAE_{p,550}$  vs.  $K^+/EC$  ratio yield a  $R^2$  of 0.23 (Figure S17c). In contrast, a much higher correlation ( $R^2=0.58$ ) was observed between  $MAE_{p,550}$  and non-BB  $MAE_{550}$  (i.e.,  $K^+/EC$  intercepts from Figure S17b). These results imply that BB is one of the contributors to the  $MAE_{p,550}$  variations, but unlikely the dominating one.

**R2-Q24.** 381: Again I am curious as to the reasoning for selection of these wavelength pairs.

**Author's Response:** The 470 and 660 nm wavelength pairs were used to represent AAE in the visible range.  $AAE_{370-880}$  ( $1.13 \pm 0.13$ ) is added as shown in Figure S8b, which is slightly higher than  $AAE_{470-660}$  ( $1.09 \pm 0.13$ ) shown in Figure S8a.

**R2-Q25.** 384-386: I do not quite follow the reasoning presented here. Why does the monthly average of 1.2 suggest variations in AAE are associated with thicker coatings than BrC contribution? To me the strongest evidence of minimal BB contribution is the lack of correlation with the  $K^+/EC$  tracer. Does AAE show any correlation with  $E_{abs}$ ? The manuscript asserts AAE is dominated by coating, so there should be a relationship based on the arguments presented.

**Author's Response:** We fully agree that the independency between AAE and  $K^+/EC$  is the main evidence that BB is not the driving force of AAE variation. The mention of monthly AAE of 1.2 here is another evidence to support this argument. The following content is revised to improve the clarity.

These results suggest that the elevated AAE observed in the PRD wintertime is unlikely to be dominated by the BB effect. Beside the independency between  $AAE_{470-660}$  and  $K^+/EC$  ratio, the measured  $AAE_{470-660}$  range also implies that BB is not the major driving force of  $AAE_{470-660}$  variations. The limited light absorption contribution from BrC in RPD region is observed in a recent study (Yuan et al., 2016), which suggest an upper limit of BrC contribution of 10% at 405 nm in the winter time using the AAE approach.

**R2-Q26.** 413-415: The refractive index of the shell also changes as it takes up water. Is this accounted for in the modeling?

**Author's Response:** The RI change due to water uptake is not considered in our Mie simulations.

**R2-Q27.** 426: This approach requires that only BC and associated coatings affect  $E_{abs}$  and AAE, but AAE can also be affected by non-BC aerosol (e.g., brown carbon or dust) that is externally mixed with BC. I am not sure how this approach can work unless there is clear evidence that there are no other light-absorbing particles or that the relative abundance does not change with season.

**Author's Response:** We add a section to discuss the caveats of the MRS method for  $E_{abs}$  determination. The MRS approach is a tracer based method. For a scenario that samples are strongly influenced by BrC (e.g. sample overall AAE>2), it is possible to determine the contribution of  $\sigma_{abs,BrC}$  if a reliable primary BrC tracer is available. We add the following content to discuss such scenarios.

The data in this study is dominated by BC absorption that did not show much influence from BrC. However, extra care should be taken if the samples exhibit substantial BrC signature (e.g. AAE>2). Such situations are equivalent to the two-source scenarios discussed in our previous paper on the MRS method (Wu and Yu, 2016) and the major findings are described below. Two types of two-source scenarios are considered: two correlated primary sources (scenario A) and two independent primary sources (scenario B). In scenario A in which both BC and primary BrC are dominated by BB, using BC as a solo tracer to calculate the primary ratio ( $MAE_p$ ) still works. In scenario B in which BC and primary BrC are independent, using BC alone to determine a single primary  $MAE_p$  could lead to a considerable bias in  $E_{abs}$  estimation. Alternatively, if a reliable primary BrC tracer is available, the corresponding  $MAE_{p,BrC}$  can be determined by MRS. With the knowledge of  $MAE_{p,BrC}$  and  $MAE_{p,BC}$ , light absorption by BC and BrC can be calculated separately and the  $E_{abs}$  can be determined using Eq. (11):

$$E_{abs} = \frac{\sigma_{abs,t}}{\sigma_{abs,p,BC} + \sigma_{abs,p,BrC}} = \frac{\sigma_{abs,t}}{MAE_{p,BC} \times EC + MAE_{p,BrC} \times BrC} \quad (11)$$

However, the implementation of Eq.11 is challenging due to the complexity in the chemical composition of BrC. For example, a recent study found that the 20 most absorbing BrC chromophores account for ~50% BrC light absorption and there is not a single compound contributing more than 10% (Lin et al., 2016), making it difficult to choose a single compound as the BrC tracer. In addition, time resolved measurement of BrC chromophores has yet to emerge. As a result, for scenario B (sample AAE>2 & primary BrC variations independent of BC), estimation of  $E_{abs}$  by MRS is not practical at this stage due to the lack of required input data. Using BC alone to determine a single primary  $MAE_p$  could lead to a considerable bias and should be avoided.

**R2-Q28.** Table 1, while convenient, could be omitted for length.

**Author's Response:** We feel that keeping table 1 in main text would be helpful for quick lookup of the abbreviations used in this study.

**R2-Q29.** Figures 2-4 do not add much to the manuscript, in my opinion, and could be removed for length.

**Author's Response:** Figure 2 is now moved to SI. As for Figure 3 and 4, they are useful to help the readers to understand how core/shell combinations can affect the variability of AAE and  $E_{abs}$  from a theoretical perspective. These two figures are used in the later discussions in section 4.5 and 4.6. Keeping these two figures will be convenient for readers to understand how observed  $E_{abs}$  and AAE can be used to infer the core/shell range from Figure 3 and 4. We believe the inclusion of the two Mie simulation figures is helpful in understand the linkage between observed AAE/ $E_{abs}$  and core/shell ranges, especially from the cross-section plots.

**R2-Q30.** Figure 7: Minor, but x-axis is year and month, not just month. Replacing with “Date” would be fine.

**Author’s Response:** Revision made.

References:

- Bond et al. (2006) Limitations in the enhancement of visible light absorption due to mixing state, *J. Geophysical Research*, doi: 10.1029/2006JD007315
- Lack and Cappa (2010) Impact of brown and clear carbon on light absorption enhancement, single scatter albedo and absorption wavelength dependence of black carbon, *Atmospheric Chemistry and Physics*, doi: 10.5194/acp-10-4207-2010
- Lack et al. (2008) Bias in filter-based light absorption measurements due to organic aerosol loading: Evidence from ambient measurements, *Aerosol Science and Technology*, 42, doi: 10.1080/02786820802389277
- Liu et al. (2015) Enhanced light absorption by mixed source black and brown carbon particles in UK winter, *Nature Communications*, doi: 10.1038/ncomms9435
- Subramanian et al. (2007) Yellow beads and missing particles: Trouble ahead for filter-based absorption measurements, *Aerosol Science and Technology*, doi:10.1080/02786820701344589

**References:**

- Ajtai, T., Filep, Á., Utry, N., Schnaiter, M., Linke, C., Bozóki, Z., Szabó, G., and Leisner, T.: Inter-comparison of optical absorption coefficients of atmospheric aerosols determined by a multi-wavelength photoacoustic spectrometer and an Aethalometer under sub-urban wintry conditions, *J. Aerosol. Sci.*, 42, 859-866, 10.1016/j.jaerosci.2011.07.008, 2011.
- Arnott, W. P., Moosmuller, H., Sheridan, P. J., Ogren, J. A., Raspert, R., Slaton, W. V., Hand, J. L., Kreidenweis, S. M., and Collett, J. L.: Photoacoustic and filter-based ambient aerosol light absorption measurements: Instrument comparisons and the role of relative humidity, *J. Geophys. Res.*, 108, 2003.
- Arnott, W. P., Hamasha, K., Moosmuller, H., Sheridan, P. J., and Ogren, J. A.: Towards aerosol light-absorption measurements with a 7-wavelength Aethalometer: Evaluation with a photoacoustic instrument and 3-wavelength nephelometer, *Aerosol. Sci. Technol.*, 39, 17-29, Doi 10.1080/027868290901972, 2005.
- Cappa, C. D., Lack, D. A., Burkholder, J. B., and Ravishankara, A. R.: Bias in Filter-Based Aerosol Light Absorption Measurements Due to Organic Aerosol Loading: Evidence from Laboratory Measurements, *Aerosol. Sci. Technol.*, 42, 1022-1032, 10.1080/02786820802389285, 2008.
- Chow, J. C., Watson, J. G., Crow, D., Lowenthal, D. H., and Merrifield, T.: Comparison of IMPROVE and NIOSH carbon measurements, *Aerosol. Sci. Technol.*, 34, 23-34, 10.1080/027868201300081923, 2001.
- Chow, J. C., Watson, J. G., Chen, L. W. A., Arnott, W. P., and Moosmuller, H.: Equivalence of elemental carbon by thermal/optical reflectance and transmittance with different temperature protocols, *Environ. Sci. Technol.*, 38, 4414-4422, 10.1021/Es034936u, 2004.

Coen, M. C., Weingartner, E., Apituley, A., Ceburnis, D., Fierz-Schmidhauser, R., Flentje, H., Henzing, J. S., Jennings, S. G., Moerman, M., Petzold, A., Schmid, O., and Baltensperger, U.: Minimizing light absorption measurement artifacts of the Aethalometer: evaluation of five correction algorithms, *Atmos. Meas. Tech.*, 3, 457-474, 10.5194/amt-3-457-2010, 2010.

Guyon, P., Graham, B., Roberts, G. C., Mayol-Bracero, O. L., Maenhaut, W., Artaxo, P., and Andreae, M. O.: Sources of optically active aerosol particles over the Amazon forest, *Atmos. Environ.*, 38, 1039-1051, 10.1016/j.atmosenv.2003.10.051, 2004.

Hoffer, A., Gelencser, A., Guyon, P., Kiss, G., Schmid, O., Frank, G. P., Artaxo, P., and Andreae, M. O.: Optical properties of humic-like substances (HULIS) in biomass-burning aerosols, *Atmos. Chem. Phys.*, 6, 3563-3570, 2006.

Kirchstetter, T. W., Novakov, T., and Hobbs, P. V.: Evidence that the spectral dependence of light absorption by aerosols is affected by organic carbon, *J. Geophys. Res.*, 109, D21208, 10.1029/2004jd004999, 2004.

Kozlov, V. S., Panchenko, M. V., Tikhomirov, A. B., Tikhomirov, B. A., and Shmargunov, V. P.: Effect of relative air humidity on photoacoustic aerosol absorption measurements in the near-ground atmospheric layer, *Atmospheric and Oceanic Optics*, 24, 487, 10.1134/s1024856011050101, 2011.

Lack, D. A. and Cappa, C. D.: Impact of brown and clear carbon on light absorption enhancement, single scatter albedo and absorption wavelength dependence of black carbon, *Atmos. Chem. Phys.*, 10, 4207-4220, DOI 10.5194/acp-10-4207-2010, 2010.

Langridge, J. M., Richardson, M. S., Lack, D. A., Brock, C. A., and Murphy, D. M.: Limitations of the Photoacoustic Technique for Aerosol Absorption Measurement at High Relative Humidity, *Aerosol. Sci. Technol.*, 47, 1163-1173, 10.1080/02786826.2013.827324, 2013.

Lee, H. M., Okuyama, K., Mizohata, A., Kim, T. O., and Koyama, H.: Fabrication of reference filter for measurements of EC (elemental carbon) and OC (organic carbon) in aerosol particles, *Aerosol. Sci. Technol.*, 41, 284-294, 2007.

Lewis, K. A., Arnott, W. P., Moosmüller, H., Chakrabarty, R. K., Carrico, C. M., Kreidenweis, S. M., Day, D. E., Malm, W. C., Laskin, A., Jimenez, J. L., Ulbrich, I. M., Huffman, J. A., Onasch, T. B., Trimborn, A., Liu, L., and Mishchenko, M. I.: Reduction in biomass burning aerosol light absorption upon humidification: roles of inorganically-induced hygroscopicity, particle collapse, and photoacoustic heat and mass transfer, *Atmos. Chem. Phys.*, 9, 8949-8966, 10.5194/acp-9-8949-2009, 2009.

Lin, P., Aiona, P. K., Li, Y., Shiraiwa, M., Laskin, J., Nizkorodov, S. A., and Laskin, A.: Molecular Characterization of Brown Carbon in Biomass Burning Aerosol Particles, *Environ. Sci. Technol.*, 50, 11815-11824, 10.1021/acs.est.6b03024, 2016.

Replacement Filter Tape for the Magee Scientific Model AE33 Aethalometer®: [http://www.mageesci.com/images/stories/docs/Magee\\_Scientific\\_Filter\\_Aethalometer\\_AE\\_Tape\\_Replacement\\_discussion.pdf](http://www.mageesci.com/images/stories/docs/Magee_Scientific_Filter_Aethalometer_AE_Tape_Replacement_discussion.pdf), 2017.

Raspet, R., Slaton, W. V., Arnott, W. P., and Moosmüller, H.: Evaporation–Condensation Effects on Resonant Photoacoustics of Volatile Aerosols, *Journal of Atmospheric and Oceanic Technology*, 20, 685-695, 10.1175/1520-0426(2003)20<685:eceorp>2.0.co;2, 2003.

Schmid, O., Artaxo, P., Arnott, W. P., Chand, D., Gatti, L. V., Frank, G. P., Hoffer, A., Schnaiter, M., and Andreae, M. O.: Spectral light absorption by ambient aerosols influenced by biomass burning in the Amazon Basin. I: Comparison and field calibration of absorption measurement techniques, *Atmos. Chem. Phys.*, 6, 3443-3462, 2006.

Subramanian, R., Roden, C. A., Boparai, P., and Bond, T. C.: Yellow beads and missing particles: Trouble ahead for filter-based absorption measurements, *Aerosol. Sci. Technol.*, 41, 630-637, Doi 10.1080/02786820701344589, 2007.

Tan, H., Liu, L., Fan, S., Li, F., Yin, Y., Cai, M., and Chan, P. W.: Aerosol optical properties and mixing state of black carbon in the Pearl River Delta, China, *Atmos. Environ.*, 131, 196-208, 10.1016/j.atmosenv.2016.02.003, 2016.

Weingartner, E., Saathoff, H., Schnaiter, M., Streit, N., Bitnar, B., and Baltensperger, U.: Absorption of light by soot particles: determination of the absorption coefficient by means of aethalometers, *J. Aerosol. Sci.*, 34, 1445-1463, 10.1016/S0021-8502(03)00359-8, 2003.

Wu, C., Ng, W. M., Huang, J., Wu, D., and Yu, J. Z.: Determination of Elemental and Organic Carbon in PM<sub>2.5</sub> in the Pearl River Delta Region: Inter-Instrument (Sunset vs. DRI Model 2001 Thermal/Optical Carbon Analyzer) and Inter-Protocol Comparisons (IMPROVE vs. ACE-Asia Protocol), *Aerosol. Sci. Technol.*, 46, 610-621, 10.1080/02786826.2011.649313, 2012.

Wu, C., Huang, X. H. H., Ng, W. M., Griffith, S. M., and Yu, J. Z.: Inter-comparison of NIOSH and IMPROVE protocols for OC and EC determination: implications for inter-protocol data conversion, *Atmos. Meas. Tech.*, 9, 4547-4560, 10.5194/amt-9-4547-2016, 2016.

Wu, C. and Yu, J. Z.: Determination of primary combustion source organic carbon-to-elemental carbon (OC/EC) ratio using ambient OC and EC measurements: secondary OC-EC correlation minimization method, *Atmos. Chem. Phys.*, 16, 5453-5465, 10.5194/acp-16-5453-2016, 2016.

Wu, D., Mao, J. T., Deng, X. J., Tie, X. X., Zhang, Y. H., Zeng, L. M., Li, F., Tan, H. B., Bi, X. Y., Huang, X. Y., Chen, J., and Deng, T.: Black carbon aerosols and their radiative properties in the Pearl River Delta region, *Sci China Ser D*, 52, 1152-1163, 10.1007/s11430-009-0115-y, 2009.

Yuan, J. F., Huang, X. F., Cao, L. M., Cui, J., Zhu, Q., Huang, C. N., Lan, Z. J., and He, L. Y.: Light absorption of brown carbon aerosol in the PRD region of China, *Atmos. Chem. Phys.*, 16, 1433-1443, 10.5194/acp-16-1433-2016, 2016.

# Quantifying black carbon light absorption enhancement by a novel statistical approach

Cheng Wu<sup>1,2</sup>, Dui Wu<sup>1,2,3</sup>, Jian Zhen Yu<sup>4,5,6</sup>

<sup>1</sup>Institute of Mass Spectrometer and Atmospheric Environment, Jinan University, Guangzhou 510632, China

<sup>2</sup>Guangdong Provincial Engineering Research Center for on-line source apportionment system of air pollution, Guangzhou 510632, China

<sup>3</sup>Institute of Tropical and Marine Meteorology, China Meteorological Administration, Guangzhou 510080, China

<sup>4</sup>Division of Environment, Hong Kong University of Science and Technology, Clear Water Bay, Hong Kong, China

<sup>5</sup>Atmospheric Research Centre, Fok Ying Tung Graduate School, Hong Kong University of Science and Technology, Nansha, China

<sup>6</sup>Department of Chemistry, Hong Kong University of Science and Technology, Clear Water Bay, Hong Kong, China

Corresponding to: Cheng Wu ([wucheng.vip@foxmail.com](mailto:wucheng.vip@foxmail.com)) and Jian Zhen Yu ([jian.yu@ust.hk](mailto:jian.yu@ust.hk))

## Abstract

Black carbon (BC) particles in the atmosphere can absorb more light when coated by non-absorbing or weakly absorbing materials during atmospheric aging, due to the lensing effect. In this study, the light absorption enhancement factor,  $E_{\text{abs}}$ , was quantified using one year's measurement of mass absorption efficiency (MAE) in the Pearl River Delta region (PRD). A new approach for calculating primary MAE ( $\text{MAE}_p$ ), the key for  $E_{\text{abs}}$  estimation, is demonstrated using the Minimum Residual Squared (MRS) method, exploring the inherent source independency between BC and its coating materials. A unique feature of  $E_{\text{abs}}$  estimation by the MRS approach is its insensitivity to systematic biases in EC and  $\sigma_{\text{abs}}$  measurements. The annual average  $E_{\text{abs}550}$  is found to be  $1.50 \pm 0.48$  ( $\pm 1$  S.D.), exhibiting a clear seasonal pattern with higher values in summer and lower in the winter. Elevated  $E_{\text{abs}}$  in the rainy summer season is likely associated with aged air masses dominating from marine origin, along with long-range transport of biomass burning influenced air masses from Southeast Asia.  $E_{\text{abs}}$



31 induced by hygroscopic growth at elevated RH could be as high as 1.3. Core-shell Mie simulations  
32 along with measured  $E_{abs}$  and Angstrom absorption exponent (AAE) constraints suggest that in the  
33 PRD, the coating materials are unlikely to be dominated by brown carbon and the coating thickness is  
34 higher in the rainy season than the dry season. A negative correlation is found between  $AAE_{470-660}$  and  
35 RH, suggesting a dominant particle size of  $D_{core} = 130$  nm and  $D_{shell}/D_{core}$  range of 2 to 4.

36

## 37 **1 Introduction**

38 Originating from incomplete combustion, black carbon (BC) is a crucial constituent of  
39 atmospheric aerosols, and is an air pollutant itself, having an adverse health impacts on humans (Suglia  
40 et al., 2008). BC has also been recognized as the third most important climate forcer due to its broad  
41 light absorbing capability across the UV-Vis-IR spectrum (IPCC, 2013). BC can alter the climate in a  
42 variety of ways, including by direct forcing (Bond et al., 2011), affecting cloud cover (Koch and Del  
43 Genio, 2010) and precipitation (Tao et al., 2012), reducing the albedo of snow and ice (Hansen and  
44 Nazarenko, 2004) and causing surface dimming (Wild, 2011). The climate effects of BC can be global  
45 or regional (Ramanathan and Carmichael, 2008). A recent study found BC can modify planetary  
46 boundary layer meteorology, and thus enhance local pollution indirectly (Ding et al., 2016). However,  
47 due to its variable optical characteristics induced during atmospheric aging, large uncertainties still  
48 exist in estimating the radiative forcing from BC. Optical properties of BC can be predicted by  
49 knowing the mass concentration, mixing state and size distribution, which collectively serve as the  
50 cornerstone for modeling the climate effect of BC. In 3D modeling studies, to conserve computational  
51 resources, the mass absorption efficiency (MAE) or mass absorption cross-section (MAC) is widely  
52 used to convert black carbon mass concentration to light absorption coefficient ( $\sigma_{abs}$ ). MAE is a  
53 quantity to describe the light absorption ability per unit EC mass:

$$54 \quad MAE (m^2 g^{-1}) = \frac{\text{absorption coefficient } \sigma_{abs} (Mm^{-1})}{EC \text{ mass concentration } (\mu g m^{-3})} \quad (1)$$

55 As a fundamental input parameter, MAE has a critical impact on BC's radiative forcing  
56 estimation in climate modeling studies. Mixing state is one of the governing factors affecting MAE.

57 Light absorption of soot particles is enhanced when coated with non-absorbing materials (Fuller et al.,  
 58 1999) or weakly absorbing materials (Lack and Cappa, 2010) during atmospheric aging. The coating  
 59 materials can focus more light onto the soot core through the lensing effect, resulting in elevated MAE  
 60 (Wang et al., 2017). Strong correlations between MAE and the number/volume fraction of coated  
 61 particles have been reported in urban areas like Tokyo (Naoe et al., 2009), Shenzhen (Lan et al., 2013)  
 62 and Xi'an (Wang et al., 2014), implying that the elevated MAE observed at these locations was mainly  
 63 due to the elevated fraction of coated of soot particles. Total absorption ( $\sigma_{abs,t}$ ) of coated particles can  
 64 be separated into two parts: primary absorption ( $\sigma_{abs,p}$ ) due to the uncoated soot core alone, and extra  
 65 absorption ( $\sigma_{abs,c}$ ) due to lensing effect of the coating (Bond et al., 2006; Jacobson, 2006; Liu et al.,  
 66 2016a) and the presence of secondarily formed brown carbon (BrC) (Lack and Cappa, 2010; Liu et al.,  
 67 2016b).

$$68 \quad \sigma_{abs,t} = \sigma_{abs,p} + \sigma_{abs,c} \quad (2)$$

69 The absorption enhancement factor ( $E_{abs}$ ) then can be defined as ratio of the total absorption and  
 70 primary absorption coefficients or the corresponding MAE values:

$$71 \quad E_{abs} = \frac{\sigma_{abs,t}}{\sigma_{abs,p}} = \frac{MAE_t}{MAE_p} \quad (3)$$

72 Where  $MAE_p$  represents the ratio of  $\sigma_{abs,p}/EC$  for uncoated soot particles, similar to the concept of  
 73 the primary OC/EC ratio in the EC tracer method:

$$74 \quad MAE_p = \frac{\sigma_{abs,p}}{EC} \quad (4)$$

75 And the MAE of coated BC can be defined as:

$$76 \quad MAE_t = \frac{\sigma_{abs,t}}{EC} \quad (5)$$

77 Thus, elevated MAE induced by coating during atmospheric aging results in an  $E_{abs}$  larger than 1.

78 Previous model studies suggest that absorption by aged soot particles can be 1.5 times greater than  
 79 fresh soot (Fuller et al., 1999; Bond et al., 2006). Laboratory studies have demonstrated that soot  
 80 particles coated with SOA (Saathoff et al., 2003; Schnaiter et al., 2005) and sulfuric acid (Zhang et al.,  
 81 2008; Khalizov et al., 2009) can increase  $E_{abs}$ . An artificial coating experiment by Shiraiwa et al. (2010)

82 found an  $E_{\text{abs}}$  of 2 for graphite particles growing in diameter from 185 to 370 nm. A laboratory study  
83 by McMeeking et al. (2014) found that in the presence of BrC, light absorption enhancement is more  
84 pronounced at the shorter wavelength. A recent chamber study coupling actual ambient air with seed  
85 BC particles implies that the timescale for  $E_{\text{abs}}$  reaching 2.4 is only 5 hours in Beijing but 18 hours in  
86 Houston (Peng et al., 2016). Field studies conducted in recent years have also substantiated enhanced  
87 light absorption in Canada (Knox et al., 2009; Chan et al., 2011), US (Lack et al., 2012b), UK (Liu et  
88 al., 2015) and Japan (Nakayama et al., 2014; Ueda et al., 2016). In contrast, field studies in California,  
89 US (Cappa et al., 2012) found a weaker light absorption enhancement (6% on average). A recent study  
90 suggests the mass ratio of non-BC content to BC particles determines the occurrence of the absorption  
91 enhancement of black-carbon particles (Liu et al., 2017).

92 Two approaches are widely used to determine  $E_{\text{abs}}$  from ambient measurements. The first approach  
93 removes the coating materials on particles physically using a thermal denuder (TD) (Lack et al., 2012a)  
94 or by aerosol filter filtration-dissolution (AFD) (Cui et al., 2016b). The TD approach is briefly  
95 discussed here. Coating materials can be removed by TD at a working temperature around 200 to  
96 300 °C (depending on the charring characteristics of aerosols at the sampling site) to measure  $\sigma_{\text{abs},p}$ ,  
97 which are cycled with measurements of  $\sigma_{\text{abs},t}$  (without passing through TD), allowing  $E_{\text{abs}}$  to be  
98 obtained from the ratio of  $\sigma_{\text{abs},t}/\sigma_{\text{abs},p}$  following Eq.3. The major advantage of the TD approach is  
99 its ability to provide highly time resolved measurements (minutes). A photo-acoustic spectrometer  
100 (PAS) is commonly used with TD for detection to satisfy its high time resolution demands. As an in-  
101 situ technique, PAS eliminates the artifacts associated with filter-based methods (Weingartner et al.,  
102 2003; Coen et al., 2010) and is often considered as the reference instrument for light absorption  
103 coefficient determination (Arnott et al., 2003; Arnott et al., 2005). One limitation of the TD approach is  
104 that a universal optimal operation temperature does not exist. If the temperature is too low, the coating  
105 cannot be fully removed, and charring can occur if the TD temperature is too high, leading to biased  
106 results. Another issue is particle loss due to TD, which can be  $\sim 20\%$  and needs to be taken into account  
107 (Ueda et al., 2016). It's also worth noting that  $\text{MAE}_p$  by the TD approach is different from the  $\text{MAE}_p$   
108 at the emission source. First, the morphology of thermally denuded BC particles (compact aggregates)  
109 is different from that of freshly emitted BC particles (chain-like aggregates). Second, most of the

110 coatings is removed with the TD denuded soot particles, but freshly emitted soot particles usually  
111 come with a thin coating of OC formed from condensation of OC vapors as the temperature drops  
112 from engine to the ambient air. As a result, the MAE<sub>p</sub> by TD approach is expected to be lower than  
113 the MAE<sub>p</sub> of emission source. In this sense, the TD approach may not be a perfect “time machine” to  
114 reverse the aging process for E<sub>abs</sub> determination.

115 The second approach is the MAE ratio method, which is also stated in Eq. 3. The key to this method  
116 is determining an appropriate MAE<sub>p</sub> that can represent the MAE from primary emissions. One  
117 approach is to adopt the reference MAE<sub>p</sub> from the literature but it may fail to represent the actual  
118 MAE<sub>p</sub> at a specific sampling site, since MAE<sub>p</sub> varies temporally and spatially. For example, MAE<sub>p</sub> of  
119 diesel soot was found to be 7.1 m<sup>2</sup>g<sup>-1</sup> at 532 nm (Adler et al., 2010). A much higher MAE<sub>p</sub> (16 m<sup>2</sup>g<sup>-1</sup>  
120 at 530 nm) was observed from natural gas flaring (Weyant et al., 2016). MAE<sub>p</sub> of biomass burning  
121 (BB) samples is highly varied due to a wide range of fuel types and combustion conditions (Reid et  
122 al., 2005; Roden et al., 2006). A range from 6.1 to 80.8 m<sup>2</sup>g<sup>-1</sup> was reported for BB MAE<sub>p</sub> at 550 nm  
123 (Pandey et al., 2016). Without the knowledge of source contributions, it is not feasible to derive a  
124 representative MAE<sub>p</sub> for E<sub>abs</sub> estimation. The other commonly used approach is to determine MAE<sub>p</sub>  
125 from the dependency of MAE on the number fraction of coated soot particles measured by SP2 (Lan  
126 et al., 2013). Since MAE (y axis) is positively correlated with the number fraction of coated soot  
127 particles (x axis), MAE<sub>p</sub> can be determined by extending the regression line to x=0. It is worth noting  
128 that this approach provides only a rough approximation of E<sub>abs</sub> since the parameter used here (coated  
129 soot particles number fraction) ignores other main drivers of light absorption enhancement (e.g.  
130 coating thickness). As a result, this approach is only valid for a period of measurements, for which  
131 coating thickness is relatively constant and the MAE variations are dominated by coated soot particles  
132 number fraction.

133 However, the high cost of the TD-PAS system and SP2 limit the field measurement of E<sub>abs</sub> around  
134 the world. In addition, long-term E<sub>abs</sub> measurements by a TD-PAS system and SP2 are not easily  
135 achieved and rarely reported. On the other hand, an Aethalometer and RT-ECOC analyzer can be  
136 effectively deployed for long term measurements and E<sub>abs</sub> estimation, at a relatively lower cost. In this  
137 study, based on one year of hourly MAE measurements (with the field carbon analyzer and

138 Aethalometer) at a suburban site in the Pearl River Delta (PRD) region of China, quantification of  
139 MAE<sub>p</sub> is demonstrated by a novel statistical approach, the Minimum R squared method (MRS) (Wu  
140 and Yu, 2016). The aim of this study is to demonstrate the capability of E<sub>abs</sub> estimation using a year-  
141 long dataset from cost-effective instrumentation. The seasonal variability of MAE, AAE and E<sub>abs</sub> in  
142 the PRD region are characterized and their dependency on air mass origin, biomass burning and RH  
143 are discussed. Abbreviations used in this study are summarized in Table 1 for a quick lookup.

## 144 **2 Ambient measurements**

145 Sampling was conducted from Feb 2012 to Jan 2013 at the suburban Nancun (NC) site (23° 0'11.82"N,  
146 113°21'18.04"E). NC, situated on the top of the highest peak (141 m ASL) in Guangzhou's Panyu  
147 district, is located at the geographic center of the PRD region, making it a representative location for  
148 average atmospheric mixing characteristics of city clusters in the PRD region. Light absorption  
149 measurements were performed by a 7-λ Aethalometer (AE-31, Magee Scientific Company, Berkeley,  
150 CA, USA). The Aethalometer was equipped with a 2.5 μm cyclone with a sampling flow rate of 4 L  
151 min<sup>-1</sup>. Weingartner's algorithm (Weingartner et al., 2003) was adopted to correct the sampling artifacts  
152 (aerosol loading, filter matrix and scattering effect) rooted in filter based method. A customized  
153 Aethalometer data processing program (Wu, 2017a) with graphical user interface was developed to  
154 perform data correction and detailed descriptions can be found in the SI (The program is available  
155 from <https://sites.google.com/site/wuchengust>). Details of the Aethalometer setup and data correction  
156 can be found in our previous paper (Wu et al., 2013).

157 EC mass concentrations were determined by a real time ECOC analyzer (Model RT-4, Sunset  
158 Laboratory Inc., Tigard, Oregon, USA). The sunset carbon analyzer was sampling on hourly cycles at  
159 a flow rate of 8 Lmin<sup>-1</sup> with a PM<sub>2.5</sub> sharp-cut cyclone inlet. For each measurement hour, the first  
160 45min were for sample collection and the remaining 15 min for thermal-optical analysis. OC is  
161 volatilized first by step-wise temperature ramping in an oxygen-free atmosphere while in the second  
162 stage EC is combusted in the presence of oxygen. Laser transmittance is applied to correct the charring  
163 artifact during the OC stage.

164 Considering a measurement **precision** of 5% for the Aethalometer (Hansen, 2005) and 24% for the  
165 RT-ECOC analyzer (Bauer et al., 2009), the propagated relative **precision** of  $E_{abs}$  ( $E_{abs,Unc}$ ) is 35%  
166 following Eq. S1&S2 in the SI. It should be noted that  $E_{abs,Unc}$  is mainly attributed to the  
167 measurement **precision** of EC by the RT-ECOC analyzer. Since the measurement **precision** of the RT-  
168 ECOC analyzer estimated by Bauer et al. (2009) is obtained from field measurement at an environment  
169 (EC below  $1 \mu\text{g m}^{-3}$ ) where EC is much lower than the present study (annual average EC  $2.63 \mu\text{g m}^{-3}$ ),  
170 the  $E_{abs,Unc}$  of 35% should be considered as an upper limit for the present study.

171 Light scattering was measured by an integrating nephelometer (Aurora-1000, Ecotech, Melbourne,  
172 Australia). Water soluble ions were measured by MARGA (The instrument for Measuring AeRosols  
173 and GAses)(ten Brink et al., 2007). **Both instruments are equipped with a PM<sub>2.5</sub> inlet to remove the**  
174 **coarse particles.**

## 175 **2.1 Uncertainties of MAE determination**

176 Two major uncertainties associated with the  $\sigma_{abs}$  and EC determination techniques should be taken  
177 into account when comparing MAE across different studies. For the  $\sigma_{abs}$  determination technique,  
178 photo-acoustic spectroscopy (PAS) is an in-situ technique free from filter based artifacts, but its  
179 application is limited by its high cost. The filter based optical transmittance method (e.g., Aethalometer  
180 and Multi Angle Absorption Photometer, MAAP) is the most widely used technique around the world,  
181 but data correction is needed to minimize the bias from artifacts due to the loading effect, matrix effect  
182 and scattering effect (Weingartner et al., 2003; Arnott et al., 2005; Schmid et al., 2006; Virkkula et al.,  
183 2007; Coen et al., 2010; Drinovec et al., 2017; Saturno et al., 2017). **Besides these artifacts, RH is also**  
184 **a source of  $\sigma_{abs}$  measurement uncertainty. Elevated RH is not only a driving force of increased  $\sigma_{abs}$**   
185 **due to the hygroscopic growth of particles, but also a factor affecting ambient  $\sigma_{abs}$  measurements.**  
186 **Previous studies found  $\sigma_{abs}$  by PAS exhibit a systematic decrease when  $\text{RH} > 70\%$  (Arnott et al., 2003;**  
187 **Kozlov et al., 2011). Water evaporation was found as the major cause for the biased PAS  $\sigma_{abs}$**   
188 **measurements under high RH (Raspert et al., 2003; Lewis et al., 2009b; Langridge et al., 2013). Filter-**  
189 **based measurements are also affected under high RH conditions. For example, Arnott et al. (2003)**  
190 **observed erratic responses by particle soot absorption photometer (PSAP) as RH changed. The main**  
191 **reason is traced to the hydrophilic cellulose membrane, which serves to reinforce the quartz filter used**  
192 **in PSAP. The fibers can swell and shrink as RH changes, causing unwanted light attenuation signal.**  
193 **The PTFE-coated glass-fiber tape has become available since 2012 for the recent model of**

194 Aethalometer to minimize the RH interference (Magee-Scientific, 2017). A study by Schmid et al.  
195 (2006) reported dependency of PSAP  $\sigma_{abs}$  on RH, but found negligible effect of RH on Aethalometer  
196 performance. It is also worth noting that RH in the Aethalometer optical chamber may be lower than  
197 the ambient RH due to the slightly elevated temperature inside the instrument. The magnitude of RH  
198 difference was found similar between different instruments: 20% for the Aethalometer (Schmid et al.,  
199 2006) and 15% for the nephelometer (Guyon et al., 2004). The RH in the Aethalometer optical chamber  
200 was not measured in this study. We expected its level to be slightly lower than the ambient RH. Cappa  
201 et al. (2008) found  $\sigma_{abs}$  measurements by PSAP and PAS maintained a high linearity even under high  
202 RH conditions (65-91%). Inter-comparison studies demonstrated that with proper corrections,  
203 Aethalometer  $\sigma_{abs}$  measurements agree well with those by PAS (Ajtai et al., 2011). During the inter-  
204 comparison study of an Aethalometer (AE-16) and a PAS in Guangzhou (Wu et al., 2009), good  
205 correlation was found ( $R^2=0.96$ ) as shown in Figure S1. These comparison results imply that the  
206 Aethalometer results are linearly correlated with PAS measurements and RH has a limited interference  
207 on Aethalometer measurements. In our study, careful corrective measures (Wu et al., 2013) are  
208 conducted for the Aethalometer  $\sigma_{abs}$  data treatment to minimize these artifacts. But such artifacts  
209 still cannot be fully eliminated.

210 For the EC determination, different thermal optical analysis (TOA) protocols can impact the  
211 measurement variability and thus MAE. As shown in Table S1, MAE for the same samples at Fresno  
212 varied from 6.1 to 9.3  $\text{m}^2 \text{g}^{-1}$ , depending on which EC analysis protocol was applied (Chow et al.,  
213 2009). Studies in the PRD found that discrepancies in measured EC by different analysis protocols  
214 could be as large as a factor of 5 (Wu et al., 2012; Wu et al., 2016a), which adds to the uncertainty for  
215 the MAE estimation. In addition, EC by TOA is also different from refractory BC (rBC) reported by  
216 the laser induced incandescence (LII) technique (e.g. single particle soot photometer, SP2). For  
217 example, two studies in Toronto (Knox et al., 2009; Chan et al., 2011) both used the PAS for  $\sigma_{abs}$   
218 measurement but different techniques for EC mass determination, resulting in very different MAE  
219 results. LII instruments are usually calibrated with a commercially available surrogate (e.g. fullerene)  
220 since direct calibration with ambient soot is not easy to achieve. Laborde et al. (2012) indicates that  
221 the incandescence response of SP2 exhibits a dependency on soot type (15% between fullerene and  
222 denuded diesel soot particles; 14% between biomass burning and denuded diesel soot particles). Due  
223 to the absence of widely accepted reference materials for EC, the uncertainties in EC determination

224 will exist in the foreseeable future. All these uncertainties, including the uncertainty of rBC mass  
225 determination by SP2, uncertainty of EC in TOA, the discrepancy between SP2 rBC and TOA EC and  
226 the discrepancy of  $\sigma_{abs}$  between filter transmission and photo-acoustic methods, can contribute to the  
227 differences in MAE listed in Table S1.

228 Systematic bias in MAE (e.g. overestimation of  $\sigma_{abs}$  and variability of EC mass by different  
229 TOA protocols) discussed above have little effect on  $E_{abs}$  estimation by MRS. As shown in Eq. 3,  $E_{abs}$   
230 is the ratio of  $MAE_t$  to  $MAE_p$  or  $\sigma_{abs,t}$  to  $\sigma_{abs,p}$ , thus most of the bias in EC mass or  $\sigma_{abs}$  is  
231 cancelled out during the  $E_{abs}$  calculation. More details are discussed in section 4.1.

## 232 **3 Methodology**

### 233 **3.1 MAE<sub>p</sub> estimation by MRS from the ambient data**

234 In this section, a new approach for MAE<sub>p</sub> estimation is introduced for  $E_{abs}$  determination, which  
235 requires the knowledge of differentiating  $\sigma_{abs,p}$  and  $\sigma_{abs,c}$  portions in  $\sigma_{abs,t}$ . The idea of  
236 decoupling  $\sigma_{abs,t}$  into  $\sigma_{abs,p}$  and  $\sigma_{abs,c}$  is conceptually similar to decoupling OC into primary OC  
237 (POC) and secondary OC (SOC) in the EC tracer method as shown in Table 2. In the EC tracer method,  
238 if  $(OC/EC)_p$  is known, POC can be determined from OC (Turpin and Huntzicker, 1991). The role of  
239 MAE<sub>p</sub> here is similar to the role of  $(OC/EC)_p$ , the primary OC/EC ratio in the EC tracer method (a  
240 comparison is given in Table 2). If MAE<sub>p</sub> (average MAE from primary emission sources) is known,  
241  $E_{abs}$  can be obtained from the ratio of  $MAE_t/MAE_p$  (Eq. 3). Therefore, the key for  $E_{abs}$  estimation is to  
242 derive an appropriate MAE<sub>p</sub>. It is worth noting that MAE<sub>p</sub> here does not represent MAE from a single  
243 or specific primary emission source, instead it reflects an average and effective MAE that has taken  
244 consideration of various primary emission sources. Thus, the MAE<sub>p</sub> is conceptually analogous to  
245  $(OC/EC)_p$  in the EC tracer method, in which the primary ratio reflects an overall ratio from primary  
246 emission sources rather than from a single primary source.

247 The Minimum R squared method (MRS) explores the inherent independency between  
248 pollutants from primary emissions (e.g., EC) and products associated with secondary formation  
249 processes (e.g., SOC,  $\sigma_{abs,c}$ ) to derive the primary ratios (e.g.,  $(OC/EC)_p$ , MAE<sub>p</sub>) in the EC tracer



250 method (Wu and Yu, 2016). When applying MRS for light absorption enhancement estimation, MRS  
251 is used to explore the inherent independency between EC and  $\sigma_{abs,c}$ , which is gained during  
252 atmospheric aging after emission. An example of MAE<sub>p</sub> estimation by MRS is shown in Figure 1.  
253 Firstly, the assumed MAE<sub>p</sub> value is varied continuously in a reasonable range (0.01 to 50 m<sup>2</sup> g<sup>-1</sup> as  
254 shown in Figure 1). Then at each hypothetical MAE<sub>p</sub>,  $\sigma_{abs,c}$  can be calculated by Eq. 6 (a combination  
255 of Eq. 2&4) using EC and  $\sigma_{abs,t}$  from ambient measurements.

$$256 \quad \sigma_{abs,c} = \sigma_{abs,t} - MAE_p \times EC \quad (6)$$

257 Accordingly, for each hypothetical MAE<sub>p</sub>, a correlation coefficient value (R<sup>2</sup>) of  $\sigma_{abs,c}$  vs.  
258 EC (i.e., R<sup>2</sup>( $\sigma_{abs,c}$ , EC)) can be obtained. The series of R<sup>2</sup>( $\sigma_{abs,c}$ , EC) values (y axis) are then plotted  
259 against the assumed MAE<sub>p</sub> values (x axis) as shown by the red curve in Figure 1. The physical meaning  
260 of this plot can be interpreted as follows. The  $\sigma_{abs,p}$  is the fraction of light absorption owing to  
261 primary emitted soot particles. As a result,  $\sigma_{abs,p}$  is well correlated with EC mass. In contrast, the  
262  $\sigma_{abs,c}$  is the fraction of light absorption gained by the lensing effect of the coating on particles after  
263 emission. The variability of  $\sigma_{abs,c}$  mainly depends on the coating thickness of the soot particles.  
264 Consequently,  $\sigma_{abs,c}$  is independent of EC mass. Since variations of EC and  $\sigma_{abs,c}$  are independent,  
265 the assumed MAE<sub>p</sub> corresponding to the minimum R<sup>2</sup>(EC,  $\sigma_{abs,c}$ ) would then represent the most  
266 statistically probable MAE<sub>p</sub> of the tested dataset.

267 A computer program (Wu, 2017b) in Igor Pro (WaveMetrics, Inc. Lake Oswego, OR, USA)  
268 was developed to facilitate MRS calculation with a user friendly graphical user interface. Another two  
269 Igor Pro based computer programs Histbox (Wu, 2017c) and Scatter Plot (Wu, 2017d) are used for  
270 generating histograms, box plots and scatter plots (with Deming regressions) presented in this study.  
271 Detailed descriptions of these computer programs can be found in the SI and the computer programs  
272 are available from <https://sites.google.com/site/wuchengust>.

### 273 **3.2 Mie simulation**

274 It can be informative to model a single soot particle using Mie theory (Bohren and Huffman,  
275 1983) and understand the theoretical range and variability of the soot particle's optical properties.

276 Three types of mixing state are widely employed for parameterization: internal mixing, external  
277 mixing and core-shell. To better represent the real situation (coating due to the aging process), a core-  
278 shell model is considered in the Mie calculation (Figure S2), which is more realistic than a volume  
279 mixture model (Bond et al., 2006). An aerosol optical closure study in the North China Plain (NCP)  
280 found that the core-shell model can provide better performance than assuming purely internal mixing  
281 and external mixing (Ma et al., 2012). A morphology study using Scanning Transmission X-ray  
282 Microscopy found that core-shell is the dominating mixing state in ambient samples (Moffet et al.,  
283 2016). It should be noted that the core-shell model assumption still has its own limitations. A single  
284 particle soot photometer (SP2) study by Sedlacek et al. (2012) reported a negative lag time between  
285 the scattering and incandescence signals in samples influenced by biomass burning, implying a near  
286 surface location of soot relative to non-absorbing materials. Near surface type mixing of soot has also  
287 been observed in Tokyo, but accounted for only 10% of total mixed soot containing particles (Moteki  
288 et al., 2014). Considering the domination of core-shell type particles in the ambient environment, the  
289 core-shell assumption in our optical model is sufficient to approximate the real situation.

290 As shown in Figure S2, fresh emitted soot particles are chain-like aggregates of small spheres  
291 (30~50 nm). After the aging process, soot particles are coated with organic and inorganic materials.  
292 Sufficient evidence has shown that the coating not only results in particle size growth, but also makes  
293 the soot core become more compact due to its collapse (Alexander et al., 2008; Zhang et al., 2008;  
294 Lewis et al., 2009a), especially under high RH conditions (Leung et al., 2017). Since the spherical like  
295 core and shell favor Mie simulation, both core and shell are considered as spheres in the Mie  
296 calculation.

297 To investigate the spectrum properties of soot particles, 11 wavelengths (370, 405, 470, 520,  
298 532, 550, 590, 660, 781, 880 and 950 nm) are considered in calculations to cover wavelengths in the  
299 most frequently used absorption measurement instruments. A refractive index (RI) of  $1.85 - 0.71i$  is  
300 adopted for soot core (Bond and Bergstrom, 2006) and 1.55 for non-absorbing coating (clear shell) in  
301 the Mie calculation for all wavelengths. Studies suggest a group of organic matter (OM), known as  
302 Brown Carbon (BrC), can absorb solar radiation at UV wavelengths (Kirchstetter et al., 2004). Thus,  
303 a BrC coating (brown shell) scenario is also considered in Mie simulation following the wavelength-

304 dependent RI suggested by Lack and Cappa (2010), which ranges from 1.55-0.059i (370 nm) to 1.55-  
305 0.0005i (950 nm). A modeling study by Bond et al. (2006) indicates that absorption amplification is  
306 not sensitive to the RI, thus the result below is not expected to be sensitive to the RI variability. Due  
307 to the spherical assumption of the BC core, a constant particle density is adopted for simplicity instead  
308 of size dependent particle density. But it is worth noting that in reality, the effective density of soot  
309 varies with particle size due to the morphology change during particle aging (Tavakoli and Olfert,  
310 2014; Dastanpour et al., 2017). Both core diameters ( $D_{\text{core}}$ ) and shell diameters ( $D_{\text{shell}}$ ) are constrained  
311 in the range of 10 ~ 3000 nm in the model simulations. The Mie calculations are implemented with a  
312 customized program (Wu, 2017e) written in Igro Pro (WaveMetrics, Inc. Lake Oswego, OR, USA)  
313 and it is available from <https://sites.google.com/site/wuchengust>. It should be noted that the core-shell  
314 type mixing state of particles is still rare in 3D atmospheric models like WRF-Chem (Matsui et al.,  
315 2013; Nordmann et al., 2014) due to computational cost limitation.

### 316 **3.2.1 Mie modeled absorption angstrom exponent (AAE)**

317 Absorption Angstrom Exponent (AAE) is a widely used parameter that describes the  
318 wavelength dependence of aerosol light absorption (Moosmuller et al., 2011), which can be written  
319 explicitly as

$$320 \quad AAE(\lambda_1, \lambda_2) = -\frac{\ln(\sigma_{abs,\lambda_1}) - \ln(\sigma_{abs,\lambda_2})}{\ln(\lambda_1) - \ln(\lambda_2)} \quad (7)$$

321 It is well known that ambient soot particles exhibit an AAE close to unity (Bond, 2001).  
322 Modeled variability in  $AAE_{470-660}$  of bare soot particles is shown in Figure S3. For soot particles with  
323  $D_{\text{core}} < 200$  nm,  $AAE_{470-660}$  is very close to 1 and decreases significantly for particles with  $D_{\text{core}} > 200$   
324 nm. Considering a typical  $D_{\text{core}}$  of fresh emitted soot particles smaller than 200 nm (Rose et al., 2006;  
325 China et al., 2013), the model results confirm the frequently observed AAE close to 1 from ambient  
326 measurements (Kirchstetter et al., 2004). Modeled variability in  $AAE_{470-660}$  of soot particles coated by  
327 non-absorbing substances (clear shell) and weakly absorbing materials (brown shell) is shown in  
328 Figure 2. Elevated AAE to ~2 is observed in the clear shell scenario (Figure 2a and 3b) for the most  
329 probable soot core particle sizes (<200 nm), which agrees well with a previous model study (Lack and

330 Cappa, 2010), implying that elevated AAE cannot be exclusively attributed to mixing with BrC. AAE  
331 elevation is more pronounced in the brown shell scenario. For soot particles with  $D_{\text{core}} < 200$  nm, brown  
332 shell  $\text{AAE}_{470-660}$  can easily reach 3 for a coating of  $D_{\text{shell}}/D_{\text{core}}=3$  (Figure 2c and 2d). These high AAE  
333 results are consistent with the previous model study (Lack and Cappa, 2010) and could partially  
334 explain the high AAE observed in measurement studies (Kirchstetter et al., 2004; Hoffer et al., 2006),  
335 since the presence of externally mixed BrC particles also contribute to the wavelength dependent light  
336 absorption.

### 337 **3.2.2 Mie modeled single scattering albedo (SSA)**

338 Variability in modeled SSA of soot particles coated by non-absorbing substances and weakly  
339 absorbing materials (e.g. BrC) is shown in Figure S4. For particles with  $D_{\text{core}} < 200$  nm and  $D_{\text{shell}}/D_{\text{core}}$   
340  $< 3$ , the SSA increases gradually (up to  $\sim 0.9$ ) with a thicker coating and behaves similarly between  
341 clear shell and brown shell scenarios.

### 342 **3.2.3 Mie modeled mass absorption efficiency (MAE)**

343 MAE is a useful indicator for soot mixing state. Variability in MAE of bare soot particles as a  
344 function of particle size at a wavelength of 550 nm is illustrated in Figure S5. The magnitude of MAE  
345 is sensitive to the soot density assumption, especially for particles  $< 200$  nm (Figure S5), but the overall  
346 trend of particle size dependency is similar between different density scenarios. MAE peaks at a  
347 particle size of 200 nm and decreases dramatically for larger particles. In our MAE calculation, a soot  
348 density of  $1.9 \text{ g cm}^{-3}$  is adopted, as suggested by Bond and Bergstrom (2006). The purpose of adopting  
349 constant density is to simplify the MAE calculation. It should be noted that the effective density of  
350 soot core is highly variable in ambient environments. For example, a study in Beijing (Zhang et al.,  
351 2016b) found a value of  $1.2 \text{ g cm}^{-3}$ . A recent chamber study found the effective density of soot can  
352 evolve from  $0.43$  to  $1.45 \text{ g cm}^{-3}$  during aging as coated by m-Xylene oxidation products (Guo et al.,  
353 2016). A study by a single-particle aerosol mass spectrometer in Guangzhou found the effective  
354 density of soot increased with particle size in the range of 400 to 1600 nm (Zhang et al., 2016a). The  
355 MAE of coated particles from different core/shell diameter combinations are shown in Figure S6. For

356 thickly coated particles, the MAE in the clear shell scenario varied as  $D_{\text{shell}}/D_{\text{core}}$  increased, but the  
357 MAE of brown shell scenario increased quasi-monotonously with  $D_{\text{shell}}/D_{\text{core}}$ .

### 358 **3.2.4 Mie modeled light absorption enhancement factor ( $E_{\text{abs}}$ )**

359  $E_{\text{abs}}$  is a better indicator for soot mixing state than MAE since it does not rely on the soot density  
360 assumption and is more suitable for comparing Mie simulations with ambient measurements. Modeled  
361 variability in  $E_{\text{abs}}$  of soot particles coated by non-absorbing substances and weakly absorbing materials  
362 (e.g. BrC) is shown in Figure 3a and 3c respectively.  $E_{\text{abs}}$  is not only sensitive to the core/shell diameter  
363 combination, but also behaves very differently on the clear and brown shell assumptions. For the clear  
364 shell scenario, when  $D_{\text{coat}}/D_{\text{core}} < 2$ ,  $E_{\text{abs}}$  does not exceed 2 for particles with different soot core sizes,  
365 but for the same  $D_{\text{coat}}/D_{\text{core}}$ , a larger soot core size yields a higher  $E_{\text{abs}}$  (Figure 3b, cross-sections of  
366 Figure 3a). If  $D_{\text{coat}}/D_{\text{core}} > 2$ ,  $E_{\text{abs}}$  could be 3 to 5 for particles with a soot core smaller than 200 nm, but  
367 for particles with a soot core larger than 200 nm, the  $E_{\text{abs}}$  is limited to  $\sim 2$  as shown in Figure 3b. For  
368 the brown shell scenario,  $E_{\text{abs}}$  increased quasi-monotonically with  $D_{\text{coat}}/D_{\text{core}}$ , and this trend is similar  
369 for different soot core sizes (Figure 3d). **The main reason behind is that in the brown shell scenario,**  
370 **both lensing effect and BrC absorption contribute to  $E_{\text{abs}}$ . As shown in Figure S7, the BrC absorption**  
371 **contribution to total  $E_{\text{abs}}$  strongly depends on coating thickness and is insensitive to soot core diameters.**  
372 **When the coating is relatively thin ( $< 5$  nm for  $\lambda@370$  nm,  $< 15$  nm for  $\lambda@550$  nm and  $< 40$  nm for**  
373  **$\lambda@880$  nm), BrC absorption contribution to the total  $E_{\text{abs}}$  is less than 20%. As the coating increases to**  
374 **a certain level ( $\sim 15$  nm for  $\lambda@370$  nm,  $\sim 35$  nm for  $\lambda@550$  nm and  $\sim 90$  nm for  $\lambda@880$  nm), BrC**  
375 **absorption contribution is comparable to the lensing effect contribution, each contributing  $\sim 50\%$  to**  
376 **the total  $E_{\text{abs}}$ . When the BrC coating is sufficiently thick ( $> 30$  nm for  $\lambda@370$  nm,  $> 90$  nm for  $\lambda@550$**   
377 **nm and  $> 110$  nm for  $\lambda@880$  nm), BrC absorption dominates the  $E_{\text{abs}}$  contribution. As a result, if BrC**  
378 **coating is indeed present in ambient samples, a strong wavelength dependent  $E_{\text{abs}}$  could be observed,**  
379 **since a BrC coating of 30 nm would be enough to induce a large amount of detectable  $E_{\text{abs}}$  in the UV**  
380 **range. Another major difference between the clear and brown shell scenarios is that, for thickly coated**  
381 **particles (e.g.  $D_{\text{coat}}/D_{\text{core}} > 2$ ), the brown shell can yield a much higher  $E_{\text{abs}}$  than the clear shell.**

382 Both primary soot size distribution and coating thickness can affect the absorption  
383 enhancement of ambient BC particles. Ambient measurements by LII found soot particle number and  
384 mass modes peaking at 110 nm and 220 nm, respectively, in the PRD (Huang et al., 2011b). A study  
385 in Shanghai found similar results (70 nm for number concentrations and 200 nm for mass  
386 concentrations)(Gong et al., 2016). Considering that the LII technique is specific for BC mass  
387 determination which is independent of BC mixing state, the size distribution reported by LII can  
388 represent the size distribution of the BC core. A study using a Micro Orifice Uniform Deposit Impactor  
389 (MOUDI) found a EC mass size distribution in the PRD exhibiting three modes peaking at ~300, ~900  
390 and ~5000 nm (Yu et al., 2010), implying a substantial coating of BC particles, and a diameter  
391 amplification of 3. BC sizing by LII is based on volume equivalent diameter (VED), while MOUDI is  
392 based on aerodynamic diameter. As a result, these two techniques do not necessarily yield similar sizes,  
393 even for the bare soot particles. The conversion between these two types of diameters involves the  
394 knowledge of particle density and morphology (drag force). A recent closure study on BC mixing state  
395 in the PRD region suggests  $\sigma_{abs}$  is dominated by coated soot particles in the range of 300~400 nm  
396 (Tan et al., 2016). Considering the dominant BC core distribution measured by SP2 (110 nm), the  
397 upper limit of  $E_{abs}$  in the PRD is roughly estimated as ~2 for the clear shell scenario (Figure 3b).

## 398 4 Results and discussions

### 399 4.1 Annual measurement statistics

400 The frequency distribution (log-normal) of  $\sigma_{abs550}$  is shown in Figure 4a, with an annual average ( $\pm 1$   
401 S.D.) of  $42.65 \pm 30.78 \text{ Mm}^{-1}$ . A log-normal distribution is also found in the EC mass concentration  
402 (Figure 4b), with an annual average of  $2.66 \pm 2.27 \mu\text{g m}^{-3}$ . Figure 4c demonstrates the yearlong  
403 frequency distribution of  $MAE_{550}$  at the NC site. The annual average  $MAE_{550}$  is  $18.75 \pm 6.16 \text{ m}^2 \text{ g}^{-1}$  and  
404 the peak ( $\pm 1$  S.D.) of the lognormal fit is  $15.70 \pm 0.22 \text{ m}^2 \text{ g}^{-1}$ . A good correlation is observed between  
405  $\sigma_{abs}$  and EC mass ( $R^2=0.92$ ) as shown in Figure 4d, and the color coding indicates a MAE dependency  
406 on RH (the RH effect will be discussed in section 4.5). Annual average  $AAE_{470-660}$  is  $1.09 \pm 0.13$  (Figure  
407 S8a), indicating that soot is the dominant absorbing substance in the PRD and the brown shell scenario

408 shown in the Mie simulation is unlikely to be important. Annual mean SSA<sub>525</sub> is 0.86±0.05 (Figure  
 409 S8c), similar to previous studies in the PRD (Jung et al., 2009; Wu et al., 2009). For comparison  
 410 purpose, MAE measured at original wavelength and MAE scaled to 550 nm following the  $\lambda^{-1}$   
 411 assumption are both shown in Table S1. The MAE comparisons discussed below are MAE at 550 nm.  
 412 MAE<sub>550</sub> by previous studies at various locations was found to cover a wide range, from 5.9 to 61.6 m<sup>2</sup>  
 413 g<sup>-1</sup>. Annual average observed MAE<sub>550</sub> at NC (18.75 m<sup>2</sup> g<sup>-1</sup>) is higher than many studies shown in Figure  
 414 5, e.g., Shenzhen (Lan et al., 2013), Beijing (Yang et al., 2009), Mexico city (Doran et al., 2007) and  
 415 Fresno (Chow et al., 2009).

416 As shown in Figure 1, the annual average MAE<sub>p,550</sub> estimated by MRS is 13 m<sup>2</sup> g<sup>-1</sup>. MAE<sub>p</sub> by  
 417 MRS represents the MAE<sub>p</sub> at the emission source, which is different from the MAE<sub>p</sub> by the TD  
 418 approach for two reasons. First, the morphology of thermally denuded BC particles (compact  
 419 aggregates) is different from that of freshly emitted BC particles (chain-like aggregates). Second, most  
 420 of the coatings are removed for TD denuded soot particles, but freshly emitted soot particles usually  
 421 come with a thin coating of OC formed from condensation of OC vapors as the temperature drops from  
 422 the flame to the ambient air. As a result, the MRS-derived MAE<sub>p</sub> is expected to be higher than the  
 423 MAE<sub>p</sub> by the TD approach. The estimated MAE<sub>p,550</sub> is higher than a previous study in Guangzhou  
 424 (7.44 m<sup>2</sup> g<sup>-1</sup>) (Andreae et al., 2008), but comparable to Xi'an (11.34 m<sup>2</sup> g<sup>-1</sup>) (Wang et al., 2014) and  
 425 Toronto (9.53~12.57 m<sup>2</sup> g<sup>-1</sup>) (Knox et al., 2009). The annual average E<sub>abs550</sub> by MRS following Eq. 3  
 426 is estimated to be 1.50±0.48 (mean ± 1 S.D.).

427 It should be noted that the E<sub>abs</sub> estimation approach demonstrated here is insensitive to the MAE  
 428 bias (e.g. overestimation of  $\sigma_{abs}$  and variability of EC mass by different TOA protocols) discussed in  
 429 section 2.1, because bias in EC mass or  $\sigma_{abs}$  is cancelled out in the E<sub>abs</sub> calculation (Eq. 3), since E<sub>abs</sub>  
 430 is the ratio of  $\sigma_{abs,t}$  to  $\sigma_{abs,p}$ . To investigate the performance of the MRS approach in response to  
 431 systematic bias in EC and  $\sigma_{abs}$ , two simple tests are conducted as shown in Figures S9 and S10 by  
 432 adding systematic biases to  $\sigma_{abs550}$  and EC. Test A represents a situation when  $\sigma_{abs}$  is overestimated  
 433 and EC is underestimated. The biased data are marked as  $\sigma'_{abs550}$  and EC' respectively, as shown  
 434 below:

$$435 \quad \sigma'_{abs550} = \sigma_{abs550} \times 2 \quad (8)$$

$$436 \quad EC' = EC \times 0.7 \quad (9)$$

437 As a result, the average  $MAE_{550}$  changed from 18.75 to 53.58  $m^2 g^{-1}$  and  $MAE_p$  changed from 13 to 37  
438  $m^2 g^{-1}$  (Figure S9). However,  $E_{abs}$  by ratio of averages remain the same (1.44).

439 In Test B, EC by different TOA protocols are compared to investigate the effect of different EC  
440 determination approaches while  $\sigma_{abs550}$  remains unchanged. EC by IMPROVE TOR protocol is  
441 calculated from NIOSH TOT EC following an empirical formula for suburban sites derived from a 3-  
442 year OCEC dataset in PRD (Wu et al., 2016a):

$$443 \quad EC_{IMP\_TOR} = 2.63 \times EC_{NSH\_TOT} + 0.05 \quad (10)$$

444 As shown in Figure S10,  $MAE_{550}$  changed from 18.75 to 7.02  $m^2 g^{-1}$  and  $MAE_p$  changed from 13 to 5  
445  $m^2 g^{-1}$ , but  $E_{abs}$  remain almost the same (1.40). Result of Test B implies that although EC is  
446 operationally defined, the discrepancy of EC between TOA protocols did not weaken the role of EC  
447 serving as a tracer for primary emissions in MRS application. These examples demonstrate that  
448 systematic biases in  $\sigma_{abs550}$  and EC have no effects on  $E_{abs}$  estimation by the MRS approach.

449 As mentioned in section 1, the definition of  $MAE_p$  by the TD approach is different from the  
450  $MAE_p$  of emission source. The TD  $MAE_p$  is expected to be slightly lower than the  $MAE_p$  of emission  
451 source. Therefore, the corresponding  $E_{abs}$  are slightly different and it should be cautioned when  
452 comparing MRS-derived  $E_{abs}$  with  $E_{abs}$  by the TD approach and Mie simulations. The  $E_{abs}$  could vary  
453 by location, depending on the coating thickness and size distribution of the primary aerosols. After  
454 undergoing atmospheric aging, the  $E_{abs}$  can be increased during transport from emission source to rural  
455 areas. The magnitude of the  $E_{abs}$  found at the NC site is comparable to other locations such as Boulder  
456 (Lack et al., 2012a) (1.38), London (Liu et al., 2015) (1.4), Shenzhen (Lan et al., 2013) (1.3), Yuncheng  
457 (Cui et al., 2016b) (2.25), Jinan (Chen et al., 2017) (2.07) and Nanjing (Cui et al., 2016a) (1.6) and is  
458 higher than studies in California (Cappa et al., 2012) (1.06), as listed in Table 3. Spectrum  $E_{abs}$  are  
459 calculated from 370 to 950 nm as shown in Figure S11.  $E_{abs}$  in the PRD exhibits a weak wavelength  
460 dependence, with slightly higher  $E_{abs}$  at the shorter wavelength (e.g.  $E_{abs370} = 1.55$ ) and is relatively  
461 lower in the IR range (e.g.  $E_{abs950} = 1.49$ ).



## 462 **4.2 Monthly characteristics of MAE, AAE and SSA**

463 Monthly variations of  $MAE_{550}$  at the NC site are shown in Figure 6a and Table S2, revealing distinct  
464 patterns of higher  $MAE_{550}$  in summer and lower in winter. On the other hand,  $AAE_{470-660}$  is lower in  
465 summer and higher in winter (Figure 6b and Table S3). Monthly  $SSA_{525}$  varied from 0.83 to 0.90  
466 without a clear seasonal pattern, as shown in Figure S12 and Table S4.  $MAE_{p,550}$  estimation for  
467 individual months is shown in Figure 6a (the purple line) and monthly  $E_{abs550}$  is calculated accordingly  
468 following Eq. 3 (Figure 6c).  $E_{abs550}$  shows clear seasonal variations, with higher values from April to  
469 August (1.52~1.97 as shown in Table S5) and relatively lower values from September to March  
470 (1.24~1.49). The highest enhancement is found in August (1.97). Factors affecting variation of  $E_{abs550}$   
471 are discussed in the following sections, including air mass origin, biomass burning and RH.

## 472 **4.3 The effect of air mass origin**

473 It's of interest to understand the seasonal variations of optical properties in the PRD. Hourly backward  
474 trajectories for the past 72 hours were calculated using NOAA's HYSPLIT (Hybrid Single Particle  
475 Lagrangian Integrated Trajectory, version 4) model (Draxier and Hess, 1998) from Feb 2012 to Jan  
476 2013 as shown in Figure S13. Cluster analysis was conducted using MeteoInfo (Wang, 2014). By  
477 examining the total spatial variance (TSV), the number of clusters was determined to be four as shown  
478 in Figure S14. Cluster 1 (C1) represents continental air masses from the north, accounting for 44.4%  
479 of total trajectories. C2 (22.8%) represents marine air masses coming from the South China Sea. C3  
480 represents air masses from the east (Taiwan island). C4 (15.8%) represents transitional air masses  
481 coming from the east coastline of China. As shown in Figure 7,  $E_{abs550}$  from C2 (1.78) is higher than  
482 other clusters (1.30 – 1.42). Further Wilcoxon-Mann-Whitney tests show that  $E_{abs550}$  from C2 is  
483 significantly higher than  $E_{abs550}$  from C1, C3 and C4 (Figure S15), implying that particles from the  
484 South China Sea cluster is likely more aged than other clusters. Air mass origin in the PRD is  
485 dominated by C2 from Apr to Aug (Figure S16a) as a result of the South China Sea monsoon in the  
486 rainy season. In contrast, the dry season is ruled by continental air masses from the north (C1) due to  
487 the influence of the northeast monsoon.  $E_{abs550}$  from C2 varied from 1.67 to 2.19, but was always

488 higher than  $E_{\text{abs}550}$  from C1 and C3 during the rainy season (Figure S16b). As a result, the domination  
489 of aged air mass from the vast ocean is one of the reasons for the much higher  $E_{\text{abs}550}$  found in the rainy  
490 season.

#### 491 **4.4 The effect of biomass burning**

492 Biomass burning (BB) and vehicular emission are the two major sources of soot particles. BC  
493 from biomass burning emission, depending on the fuel type and burning condition, may have a higher  
494 OC/EC ratio and a thicker coating, resulting in a higher MAE than vehicular emission (Shen et al.,  
495 2013; Cheng et al., 2016). In this study, the influence of BB on optical properties is investigated using  
496 the  $K^+/\text{EC}$  ratio as a BB indicator. As shown in Figure 8,  $\text{MAE}_{550}$  is positively correlated with the  
497  $K^+/\text{EC}$  ratio, which exhibits a clear seasonal pattern that is higher in the rainy season and lower in the  
498 dry season (Figure S17a). Southeast Asia has the highest fire emission density globally due to the high  
499 biofuel consumption along with frequent fire activity in this region (Aouizerats et al., 2015), making  
500 Southeast Asia a large contributor to BC emissions (Jason Blake, 2014). During the rainy season when  
501 oceanic wind prevails, BC from BB emission in Southeast Asia can reach PRD through long range  
502 transport (LRT), resulting in an elevated  $K^+/\text{EC}$  ratio and  $\text{MAE}_{550}$ . The Deming regression intercept  
503 (11.89) in Figure 8 represents the MAE without the BB effect. This non-BB  $\text{MAE}_{550}$  ( $11.89 \text{ m}^2 \text{ g}^{-1}$ ) is  
504 only slightly lower than  $\text{MAE}_{\text{p},550}$  ( $13 \text{ m}^2 \text{ g}^{-1}$ ) obtained in section 4.3, implying that a large fraction of  
505  $\text{MAE}_{\text{p},550}$  could not be explained by the BB source. Additional evidence was obtained through  
506 examining regression relationships of  $\text{MAE}_{\text{p},550}$  with  $K^+/\text{EC}$  month-by-month (Figure S17b).  
507 Correlation of monthly  $\text{MAE}_{\text{p},550}$  vs.  $K^+/\text{EC}$  ratio yield a  $R^2$  of 0.23 (Figure S17c). In contrast, a much  
508 higher correlation ( $R^2=0.58$ ) was observed (Figure S17d) between  $\text{MAE}_{\text{p},550}$  and non-BB  $\text{MAE}_{550}$  (i.e.,  
509  $K^+/\text{EC}$  intercepts from Figure S17b). These results imply that BB is one of the contributors to the  
510  $\text{MAE}_{\text{p},550}$  variations, but unlikely the dominating one.

511 Many studies have found that BB influenced samples exhibit elevated AAE due to the presence  
512 of wavelength dependent light absorbing substances like BrC and HUmic-Like Substances (HULIS)  
513 (Kirchstetter et al., 2004; Hoffer et al., 2006; Sandradewi et al., 2008; Herich et al., 2011; Pokhrel et  
514 al., 2017). It is of interest to investigate whether elevated AAE observed in the PRD during the dry

515 season is associated with BB influence. As shown in Figure S18,  $AAE_{370-470}$  and  $AAE_{470-660}$  did not  
516 correlate with the BB indicator,  $K^+/EC$  ratio. These results suggest that the elevated AAE observed in  
517 the PRD wintertime is unlikely to be dominated by the BB effect. Beside the independency between  
518  $AAE_{470-660}$  and  $K^+/EC$  ratio, the measured  $AAE_{470-660}$  range also implies that BB is not the major  
519 driving force of  $AAE_{470-660}$  variations. The limited light absorption contribution from BrC in RPD  
520 region is observed in a recent study (Yuan et al., 2016) , which suggest an upper limit of BrC  
521 contribution of 10% at 405 nm in the winter time using the AAE approach. As discussed in our Mie  
522 simulation (section 3.1) and a previous study (Lack and Cappa, 2010), coating of non-absorbing  
523 materials onto soot particles can increase AAE up to 2. Since the monthly average  $AAE_{470-660}$  in  
524 wintertime did not exceed 1.2 (Table S3), the variations of  $AAE_{470-660}$  in the PRD are more likely  
525 associated with coatings rather than the contribution of BrC. The results also imply that attempts on  
526 BrC absorption attribution for the PRD dataset presented in this study could be risky, considering that  
527 elevation of AAE is actually dominated by coating (Lack and Langridge, 2013).

#### 528 **4.5 The effect of relative humidity (RH) on optical properties**

529 Soot particles are relatively hydrophobic when freshly emitted, but tend to gain hygroscopicity  
530 during atmospheric aging. Hygroscopic growth of coated laboratory generated model BC was reported  
531 by McMeeking et al. (2011). Growth of ambient BC particle size by a factor of 1.4-1.6 under high RH  
532 has been observed in a UK study (Liu et al., 2013). Located in the subtropical zone, RH plays an  
533 important role on aerosol optical properties in the PRD region. The yearlong measurements at the NC  
534 site provide a unique opportunity to investigate the effect of RH on aerosol optical properties, since  
535 most existing ad hoc studies in the PRD only last for months. Liquid water content (LWC) was  
536 calculated using the E-AIM (model 2) thermodynamic model (Clegg et al., 1998). As shown in Figure  
537 S19, LWC on average accounted for a significant fraction (44%) of non-EC  $PM_{2.5}$  mass, making it an  
538 important component of  $PM_{2.5}$  mass and due to high RH in the PRD. Previously, hygroscopic growth  
539 was only considered for particle scattering in the IMPROVE formula for chemically resolved light  
540 extinction budget studies. In this study  $f(RH)$  of MAE was obtained from yearlong measurements as  
541 shown in Figure 9a for  $RH = 30 \sim 100\%$  and color coded for LWC. It clearly shows that  $MAE_{550}$

542 measured in NC is positively correlated with RH and the enhancement can be fitted by a polynomial  
543 equation. When RH is close to 100%, the LWC can account for 70% of PM<sub>2.5</sub> mass. The maximum  
544  $f(\text{RH})$  can reach 1.3, which is higher than the value found in Beijing (1.2) (Wu et al., 2016b), **but lower**  
545 **than a numerical study (1.35) (Nessler et al., 2005)**. These results reveal that a large contribution of  
546  $E_{\text{abs}}$  is coming from high LWC under high RH in the PRD region. Because RH has a clear diurnal  
547 pattern, it can affect the diurnal pattern of  $E_{\text{abs}}$  in the PRD. Since the RH effect on  $E_{\text{abs}}$  is rarely  
548 considered in existing climate models, the inclusion of RH effect can reduce the uncertainty for  
549 assessing BC's climate effect.

550 The  $\text{AAE}_{470-660}$  dependency on RH is shown in Figure 9b. When RH is low (e.g. 30%), the  
551  $\text{AAE}_{470-660}$  is around 1.25 and decreases to 1.10 as RH increases to 50%.  $\text{AAE}_{470-660}$  remains around  
552 1.12 when RH is 50-70%. Then  $\text{AAE}_{470-660}$  decreases again when RH is higher than 70% and can reach  
553 1 when RH is close to 100%. Since a higher RH results in hygroscopic growth and larger particle  
554 diameters, the negative correlation between  $\text{AAE}_{470-660}$  and RH provides a clue on soot particles'  
555 primary diameter and mixing state. As shown in the Mie simulation in Figure 2b, for a particle with  
556  $D_{\text{core}}$  of 130 nm and  $D_{\text{shell}}/D_{\text{core}}$  of 2 to 4,  $\text{AAE}_{470-660}$  decreases as the coating increases, and the decrease  
557 tapers off when  $D_{\text{shell}}/D_{\text{core}} = 3$ . The  $D_{\text{core}}$  obtained here (130nm) is comparable with  $D_{\text{core}}$  obtained  
558 from SP2 measurements (110nm) in the PRD (Huang et al., 2011a).

## 559 **4.6 Implications for mixing state**

560 Quantitative direct measurements of BC mixing state and coating thickness are still challenging.  
561 SP2 can estimate the coating thickness using a lag-time approach or a Mie calculation approach can  
562 be employed, but both methods have a limited range in coating thickness and uncertainties arise from  
563 the assumptions made during the retrieval. For example, recent studies found that the mass equivalent  
564 diameter of soot core measured by SP2 could be underestimated due to density assumptions (Zhang et  
565 al., 2016b). Although size distribution measurement is not available in this study, clues of mixing state  
566 still can be derived from bulk measurements of optical properties. As discussed in section 4.4.1,  
567 elevated  $E_{\text{abs}550}$  observed in the rainy season is associated with aged air masses from a marine origin.  
568 To probe the possible mixing state difference between dry and rainy season,  $E_{\text{abs}550}$ ,  $\text{SSA}_{525}$  and

569 AAE<sub>470-660</sub> are used to narrow down the possible core-shell size range as shown in Figure S20. Monthly  
 570 averages with one standard deviation of AAE<sub>470-660</sub>, SSA<sub>525</sub> and E<sub>abs550</sub> are used as constraints to extract  
 571 the intersecting core-shell size range from Figure 2a, Figure S4 and Figure 3a. **January and August**  
 572 **data are used to represent two different scenarios: elevated AAE<sub>470-660</sub> (1.19±0.11) with lower E<sub>abs550</sub>**  
 573 **(1.31±0.32) in dry season and low AAE<sub>470-660</sub> (1.04±0.09) with elevated E<sub>abs550</sub> (1.97±0.71) in rainy**  
 574 **season.** The results show that January and August have a very different core-shell size range: in  
 575 January, the core and shell range are 100 ~ 160 nm and 120 ~ 250 nm, respectively; in August, the  
 576 core and shell range are 120 ~ 165 nm and 170 ~ 430 nm, respectively. This confirms again that the  
 577 soot particles in the rainy season are likely to have a thicker coating than in the dry season.

## 578 **5 Caveats of the MRS method in its applications to ambient data**

579 The data in this study is dominated by BC absorption that did not show much influence from  
 580 BrC. However, extra care should be taken if the samples exhibit substantial BrC signature (e.g.  
 581 AAE>2). Such situations are equivalent to the two-source scenarios discussed in our previous paper  
 582 on the MRS method (Wu and Yu, 2016) and the major findings are described below. Two types of  
 583 two-source scenarios are considered: two correlated primary sources (scenario A) and two independent  
 584 primary sources (scenario B). In scenario A in which both BC and primary BrC are dominated by BB,  
 585 using BC as a solo tracer to calculate the primary ratio (MAE<sub>p</sub>) still works. In scenario B in which BC  
 586 and primary BrC are independent, using BC alone to determine a single primary MAE<sub>p</sub> could lead to  
 587 a considerable bias in E<sub>abs</sub> estimation. Alternatively, if a reliable primary BrC tracer is available, the  
 588 corresponding MAE<sub>p,BrC</sub> can be determined by MRS. With the knowledge of MAE<sub>p,BrC</sub> and MAE<sub>p,BC</sub>, light  
 589 absorption by BC and BrC can be calculated separately and the E<sub>abs</sub> can be determined using Eq. (11) :

$$590 \quad E_{abs} = \frac{\sigma_{abs,t}}{\sigma_{abs,p,BC} + \sigma_{abs,p,BrC}} = \frac{\sigma_{abs,t}}{MAE_{p,BC} \times EC + MAE_{p,BrC} \times BrC} \quad (11)$$

591 However, the implementation of Eq.11 is challenging due to the complexity in the chemical  
 592 composition of BrC. For example, a recent study found that the 20 most absorbing BrC chromophores  
 593 account for ~50% BrC light absorption and there is not a single compound contributing more than 10%  
 594 (Lin et al., 2016), making it difficult to choose a single compound as the BrC tracer. In addition, time

595 resolved measurement of BrC chromophores has yet to emerge. As a result, for scenario B (sample  
596 AAE>2 & primary BrC variations independent of BC), estimation of  $E_{abs}$  by MRS is not practical at  
597 this stage due to the lack of required input data. Using BC alone to determine a single primary MAE<sub>p</sub>  
598 could lead to a considerable bias and should be avoided.

## 599 **6 Conclusions**

600 In this study, a novel statistical approach is proposed and its application on ambient data is  
601 demonstrated using one-year hourly OC and EC data coupled with Aethalometer measurements.  
602 Unlike conventional  $E_{abs}$  determination approaches that require expensive instrumentation (e.g. TD-  
603 PAS, VTDMA, SP2), this new approach employs widely deployed instruments (field carbon analyzer  
604 and Aethalometer). The key of this new approach involves calculating MAE<sub>p</sub> by the Minimum R  
605 Squared (MRS) method (Wu and Yu, 2016). It is found that  $E_{abs}$  estimation by MRS is insensitive to  
606 systematic biases in EC and  $\sigma_{abs}$  measurements. The annual average MAE<sub>p,550</sub> estimated by MRS is  
607  $13 \text{ m}^2 \text{ g}^{-1}$  and annual average MAE<sub>550</sub> is  $18.75 \pm 6.16 \text{ m}^2 \text{ g}^{-1}$ , suggesting an annual average enhancement  
608 factor ( $E_{abs550}$ ) of  $1.50 \pm 0.48$ . This value is within the upper limit of  $E_{abs}$  ( $\sim 2$ ) by core-shell Mie  
609 simulations considering the typical soot size distribution and coating thickness in the PRD.

610 Both MAE<sub>p,550</sub> and  $E_{abs}$  show distinct seasonal variations, implying the complexity of soot  
611 particle mixing state variations in this region. The elevated summertime  $E_{abs550}$  in the PRD is found to  
612 be associated with the domination of aged air masses from the South China Sea, along with the long-  
613 range transport of biomass burning influenced air masses from Southeast Asia. Hygroscopic growth  
614 with elevated RH contributes to  $E_{abs}$  as well, which could be as high as 1.3. A negative correlation is  
615 found between AAE<sub>470-660</sub> and RH, suggesting a dominant particle size with a  $D_{core}$  of 130 nm and  
616  $D_{shell}/D_{core}$  range of 2 to 4. Core-shell size ranges narrowed down by  $E_{abs550}$  and AAE<sub>470-660</sub> constraints  
617 suggest that soot particles in the rainy season are likely to have thicker coatings than in the dry season.

## 618 **Data availability**

619 OC, EC, inorganic ions and  $\sigma_{abs}$  data used in this study are available from corresponding authors  
620 upon request.

621

622 **Acknowledgements**

623 This work is supported by the National Natural Science Foundation of China (41605002, 41475004).

624 We gratefully acknowledge the Fok Ying Tung Foundation for funding to the Atmospheric Research  
625 Center at HKUST Fok Ying Tung Graduate School. The authors thank Jingxiang Huang of Fok Ying  
626 Tung Graduate School for the assistance in OCEC analyzer maintenance. The authors are also grateful  
627 to Dr. Stephen M Griffith and Dr. Yongjie Li for the helpful comments. The authors gratefully  
628 acknowledge the NOAA Air Resources Laboratory (ARL) for the provision of the HYSPLIT transport  
629 and dispersion model used in this publication.

630 **References**

631

- 632 Adler, G., Riziq, A. A., Erlick, C., and Rudich, Y.: Effect of intrinsic organic carbon on the optical  
633 properties of fresh diesel soot, *Proceedings of the National Academy of Sciences*, 107, 6699-6704, doi:  
634 10.1073/pnas.0903311106, 2010.
- 635 Ajtai, T., Filep, Á., Utry, N., Schnaiter, M., Linke, C., Bozóki, Z., Szabó, G., and Leisner, T.: Inter-  
636 comparison of optical absorption coefficients of atmospheric aerosols determined by a multi-  
637 wavelength photoacoustic spectrometer and an Aethalometer under sub-urban wintry conditions, *J.*  
638 *Aerosol. Sci.*, 42, 859-866, doi: 10.1016/j.jaerosci.2011.07.008, 2011.
- 639 Alexander, D. T. L., Crozier, P. A., and Anderson, J. R.: Brown carbon spheres in East Asian outflow  
640 and their optical properties, *Science*, 321, 833-836, 2008.
- 641 Andreae, M. O., Schmid, O., Yang, H., Chand, D., Yu, J. Z., Zeng, L. M., and Zhang, Y. H.: Optical  
642 properties and chemical composition of the atmospheric aerosol in urban Guangzhou, China, *Atmos.*  
643 *Environ.*, 42, 6335-6350, doi: 10.1016/j.atmosenv.2008.01.030, 2008.
- 644 Aouizerats, B., van der Werf, G. R., Balasubramanian, R., and Betha, R.: Importance of transboundary  
645 transport of biomass burning emissions to regional air quality in Southeast Asia during a high fire  
646 event, *Atmos. Chem. Phys.*, 15, 363-373, doi: 10.5194/acp-15-363-2015, 2015.
- 647 Arnott, W. P., Moosmuller, H., Sheridan, P. J., Ogren, J. A., Raspet, R., Slaton, W. V., Hand, J. L.,  
648 Kreidenweis, S. M., and Collett, J. L.: Photoacoustic and filter-based ambient aerosol light absorption  
649 measurements: Instrument comparisons and the role of relative humidity, *J. Geophys. Res.*, 108, 2003.
- 650 Arnott, W. P., Hamasha, K., Moosmuller, H., Sheridan, P. J., and Ogren, J. A.: Towards aerosol light-  
651 absorption measurements with a 7-wavelength Aethalometer: Evaluation with a photoacoustic  
652 instrument and 3-wavelength nephelometer, *Aerosol. Sci. Technol.*, 39, 17-29, doi: Doi  
653 10.1080/027868290901972, 2005.
- 654 Bauer, J. J., Yu, X.-Y., Cary, R., Laulainen, N., and Berkowitz, C.: Characterization of the sunset semi-  
655 continuous carbon aerosol analyzer, *J. Air Waste Manage. Assoc.*, 59, 826-833, doi: 10.3155/1047-  
656 3289.59.7.826, 2009.
- 657 Bohren, C. F. and Huffman, D. R.: *Absorption and scattering of light by small particles*, Wiley, New  
658 York, xiv, 530 p. pp., 1983.
- 659 Bond, T. C.: Spectral dependence of visible light absorption by carbonaceous particles emitted from  
660 coal combustion, *Geophys. Res. Lett.*, 28, 4075-4078, doi: Doi 10.1029/2001gl013652, 2001.
- 661 Bond, T. C. and Bergstrom, R. W.: Light absorption by carbonaceous particles: An investigative  
662 review, *Aerosol. Sci. Technol.*, 40, 27-67, doi: Doi 10.1080/02786820500421521, 2006.
- 663 Bond, T. C., Habib, G., and Bergstrom, R. W.: Limitations in the enhancement of visible light  
664 absorption due to mixing state, *J. Geophys. Res.*, 111, -, 2006.
- 665 Bond, T. C., Zarzycki, C., Flanner, M. G., and Koch, D. M.: Quantifying immediate radiative forcing  
666 by black carbon and organic matter with the Specific Forcing Pulse, *Atmos. Chem. Phys.*, 11, 1505-  
667 1525, doi: 10.5194/acp-11-1505-2011, 2011.
- 668 Cappa, C. D., Lack, D. A., Burkholder, J. B., and Ravishankara, A. R.: Bias in Filter-Based Aerosol  
669 Light Absorption Measurements Due to Organic Aerosol Loading: Evidence from Laboratory  
670 Measurements, *Aerosol. Sci. Technol.*, 42, 1022-1032, doi: 10.1080/02786820802389285, 2008.



671 Cappa, C. D., Onasch, T. B., Massoli, P., Worsnop, D. R., Bates, T. S., Cross, E. S., Davidovits, P.,  
672 Hakala, J., Hayden, K. L., Jobson, B. T., Kolesar, K. R., Lack, D. A., Lerner, B. M., Li, S.-M., Mellon,  
673 D., Nuaaman, I., Olfert, J. S., Petäjä, T., Quinn, P. K., Song, C., Subramanian, R., Williams, E. J., and  
674 Zaveri, R. A.: Radiative Absorption Enhancements Due to the Mixing State of Atmospheric Black  
675 Carbon, *Science*, 337, 1078-1081, doi: 10.1126/science.1223447, 2012.

676 Chan, T. W., Brook, J. R., Smallwood, G. J., and Lu, G.: Time-resolved measurements of black carbon  
677 light absorption enhancement in urban and near-urban locations of southern Ontario, Canada, *Atmos.*  
678 *Chem. Phys.*, 11, 10407-10432, 2011.

679 Chen, B., Bai, Z., Cui, X., Chen, J., Andersson, A., and Gustafsson, Ö.: Light absorption enhancement  
680 of black carbon from urban haze in Northern China winter, *Environ Pollut*, 221, 418-426, doi:  
681 10.1016/j.envpol.2016.12.004, 2017.

682 Cheng, Y., Engling, G., Moosmüller, H., Arnott, W. P., Chen, L. W. A., Wold, C. E., Hao, W. M., and  
683 He, K.-b.: Light absorption by biomass burning source emissions, *Atmos. Environ.*, 127, 347-354, doi:  
684 10.1016/j.atmosenv.2015.12.045, 2016.

685 China, S., Mazzoleni, C., Gorkowski, K., Aiken, A. C., and Dubey, M. K.: Morphology and mixing  
686 state of individual freshly emitted wildfire carbonaceous particles, *Nat Commun*, 4, doi:  
687 10.1038/ncomms3122, 2013.

688 Chow, J. C., Watson, J. G., Doraiswamy, P., Chen, L. W. A., Sodeman, D. A., Lowenthal, D. H., Park,  
689 K., Arnott, W. P., and Motallebi, N.: Aerosol light absorption, black carbon, and elemental carbon at  
690 the Fresno Supersite, California, *Atmos Res*, 93, 874-887, doi: DOI 10.1016/j.atmosres.2009.04.010,  
691 2009.

692 Clegg, S. L., Brimblecombe, P., and Wexler, A. S.: Thermodynamic Model of the System  
693  $H^+ - NH_4^+ - SO_4^{2-} - NO_3^- - H_2O$  at Tropospheric Temperatures, *The Journal of Physical Chemistry A*,  
694 102, 2137-2154, doi: 10.1021/jp973042r, 1998.

695 Coen, M. C., Weingartner, E., Apituley, A., Ceburnis, D., Fierz-Schmidhauser, R., Flentje, H.,  
696 Henzing, J. S., Jennings, S. G., Moerman, M., Petzold, A., Schmid, O., and Baltensperger, U.:  
697 Minimizing light absorption measurement artifacts of the Aethalometer: evaluation of five correction  
698 algorithms, *Atmos. Meas. Tech.*, 3, 457-474, doi: 10.5194/amt-3-457-2010, 2010.

699 Cui, F., Chen, M., Ma, Y., Zheng, J., Zhou, Y., Li, S., Qi, L., and Wang, L.: An intensive study on  
700 aerosol optical properties and affecting factors in Nanjing, China, *Journal of Environmental Sciences*,  
701 40, 35-43, doi: 10.1016/j.jes.2015.08.017, 2016a.

702 Cui, X., Wang, X., Yang, L., Chen, B., Chen, J., Andersson, A., and Gustafsson, Ö.: Radiative  
703 absorption enhancement from coatings on black carbon aerosols, *Sci.Total.Environ.*, 551, 51-56, doi:  
704 10.1016/j.scitotenv.2016.02.026, 2016b.

705 Dastanpour, R., Momenimovahed, A., Thomson, K., Olfert, J., and Rogak, S.: Variation of the optical  
706 properties of soot as a function of particle mass, *Carbon*, 124, 201-211, doi:  
707 10.1016/j.carbon.2017.07.005, 2017.

708 Ding, A. J., Huang, X., Nie, W., Sun, J. N., Kerminen, V. M., Petäjä, T., Su, H., Cheng, Y. F., Yang,  
709 X. Q., Wang, M. H., Chi, X. G., Wang, J. P., Virkkula, A., Guo, W. D., Yuan, J., Wang, S. Y., Zhang,  
710 R. J., Wu, Y. F., Song, Y., Zhu, T., Zilitinkevich, S., Kulmala, M., and Fu, C. B.: Enhanced haze  
711 pollution by black carbon in megacities in China, *Geophys. Res. Lett.*, 43, 2873-2879, doi:  
712 10.1002/2016GL067745, 2016.

713 Doran, J. C., Barnard, J. C., Arnott, W. P., Cary, R., Coulter, R., Fast, J. D., Kassianov, E. I., Kleinman,  
714 L., Laulainen, N. S., Martin, T., Paredes-Miranda, G., Pekour, M. S., Shaw, W. J., Smith, D. F.,  
715 Springston, S. R., and Yu, X. Y.: The T1-T2 study: evolution of aerosol properties downwind of  
716 Mexico City, *Atmos. Chem. Phys.*, 7, 1585-1598, doi: 10.5194/acp-7-1585-2007, 2007.

717 Draxier, R. R. and Hess, G. D.: An overview of the HYSPLIT\_4 modelling system for trajectories,  
718 dispersion and deposition, *Aust Meteorol Mag*, 47, 295-308, 1998.

719 Drinovec, L., Gregorič, A., Zotter, P., Wolf, R., Bruns, E. A., Prévôt, A. S. H., Petit, J. E., Favez, O.,  
720 Sciare, J., Arnold, I. J., Chakrabarty, R. K., Moosmüller, H., Filep, A., and Močnik, G.: The filter-  
721 loading effect by ambient aerosols in filter absorption photometers depends on the coating of the  
722 sampled particles, *Atmos. Meas. Tech.*, 10, 1043-1059, doi: 10.5194/amt-10-1043-2017, 2017.

723 Fuller, K. A., Malm, W. C., and Kreidenweis, S. M.: Effects of mixing on extinction by carbonaceous  
724 particles, *J. Geophys. Res.*, 104, 15941-15954, 1999.

725 Gong, X., Zhang, C., Chen, H., Nizkorodov, S. A., Chen, J., and Yang, X.: Size distribution and mixing  
726 state of black carbon particles during a heavy air pollution episode in Shanghai, *Atmos. Chem. Phys.*,  
727 16, 5399-5411, doi: 10.5194/acp-16-5399-2016, 2016.

728 Guo, S., Hu, M., Lin, Y., Gomez-Hernandez, M., Zamora, M. L., Peng, J., Collins, D. R., and Zhang,  
729 R.: OH-Initiated Oxidation of m-Xylene on Black Carbon Aging, *Environ. Sci. Technol.*, doi:  
730 10.1021/acs.est.6b01272, 2016.

731 Guyon, P., Graham, B., Roberts, G. C., Mayol-Bracero, O. L., Maenhaut, W., Artaxo, P., and Andreae,  
732 M. O.: Sources of optically active aerosol particles over the Amazon forest, *Atmos. Environ.*, 38, 1039-  
733 1051, doi: 10.1016/j.atmosenv.2003.10.051, 2004.

734 Hansen, A. D. A.: *The Aethalometer Manual*, Berkeley, California, USA, Magee Scientific, 2005.

735 Hansen, J. and Nazarenko, L.: Soot climate forcing via snow and ice albedos, *P Natl Acad Sci USA*,  
736 101, 423-428, doi: DOI 10.1073/pnas.2237157100, 2004.

737 Herich, H., Hueglin, C., and Buchmann, B.: A 2.5 year's source apportionment study of black carbon  
738 from wood burning and fossil fuel combustion at urban and rural sites in Switzerland, *Atmos. Meas.*  
739 *Tech.*, 4, 1409-1420, doi: DOI 10.5194/amt-4-1409-2011, 2011.

740 Hoffer, A., Gelencser, A., Guyon, P., Kiss, G., Schmid, O., Frank, G. P., Artaxo, P., and Andreae, M.  
741 O.: Optical properties of humic-like substances (HULIS) in biomass-burning aerosols, *Atmos. Chem.*  
742 *Phys.*, 6, 3563-3570, 2006.

743 Huang, X. F., Gao, R. S., Schwarz, J. P., He, L. Y., Fahey, D. W., Watts, L. A., McComiskey, A.,  
744 Cooper, O. R., Sun, T. L., Zeng, L. W., Hu, M., and Zhang, Y. H.: Black carbon measurements in the  
745 Pearl River Delta region of China, *J. Geophys. Res.*, 116, D12208, doi: 10.1029/2010jd014933, 2011a.

746 Huang, X. F., He, L. Y., Hu, M., Canagaratna, M. R., Kroll, J. H., Ng, N. L., Zhang, Y. H., Lin, Y.,  
747 Xue, L., Sun, T. L., Liu, X. G., Shao, M., Jayne, J. T., and Worsnop, D. R.: Characterization of  
748 submicron aerosols at a rural site in Pearl River Delta of China using an Aerodyne High-Resolution  
749 Aerosol Mass Spectrometer, *Atmos. Chem. Phys.*, 11, 1865-1877, doi: 10.5194/acp-11-1865-2011,  
750 2011b.

751 IPCC: *Climate change 2013 : the physical science basis : Working Group I contribution to the Fifth*  
752 *Assessment Report of the Intergovernmental Panel on Climate Change*, xi, 1535 pages. pp., 2013.

753 Jacobson, M. Z.: Effects of externally-through-internally-mixed soot inclusions within clouds and  
754 precipitation on global climate, *J Phys Chem A*, 110, 6860-6873, 2006.

755 Jason Blake, C.: Quantifying the occurrence and magnitude of the Southeast Asian fire climatology,  
756 *Environmental Research Letters*, 9, 114018, 2014.

757 Jung, J., Lee, H., Kim, Y. J., Liu, X., Zhang, Y., Gu, J., and Fan, S.: Aerosol chemistry and the effect  
758 of aerosol water content on visibility impairment and radiative forcing in Guangzhou during the 2006  
759 Pearl River Delta campaign, *Journal of Environmental Management*, 90, 3231-3244, doi:  
760 10.1016/j.jenvman.2009.04.021, 2009.

761 Khalizov, A. F., Xue, H. X., Wang, L., Zheng, J., and Zhang, R. Y.: Enhanced Light Absorption and  
762 Scattering by Carbon Soot Aerosol Internally Mixed with Sulfuric Acid, *J Phys Chem A*, 113, 1066-  
763 1074, 2009.

764 Kirchstetter, T. W., Novakov, T., and Hobbs, P. V.: Evidence that the spectral dependence of light  
765 absorption by aerosols is affected by organic carbon, *J. Geophys. Res.*, 109, D21208, doi:  
766 10.1029/2004jd004999, 2004.

767 Knox, A., Evans, G. J., Brook, J. R., Yao, X., Jeong, C. H., Godri, K. J., Sabaliauskas, K., and Slowik,  
768 J. G.: Mass Absorption Cross-Section of Ambient Black Carbon Aerosol in Relation to Chemical Age,  
769 *Aerosol. Sci. Technol.*, 43, 522-532, doi: Doi 10.1080/02786820902777207, 2009.

770 Koch, D. and Del Genio, A.: Black carbon semi-direct effects on cloud cover: review and synthesis,  
771 *Atmos. Chem. Phys.*, 10, 7685-7696, 2010.

772 Kozlov, V. S., Panchenko, M. V., Tikhomirov, A. B., Tikhomirov, B. A., and Shmargunov, V. P.:  
773 Effect of relative air humidity on photoacoustic aerosol absorption measurements in the near-ground  
774 atmospheric layer, *Atmospheric and Oceanic Optics*, 24, 487, doi: 10.1134/s1024856011050101, 2011.

775 Laborde, M., Mertes, P., Zieger, P., Dommen, J., Baltensperger, U., and Gysel, M.: Sensitivity of the  
776 Single Particle Soot Photometer to different black carbon types, *Atmos. Meas. Tech.*, 5, 1031-1043,  
777 2012.

778 Lack, D. A. and Cappa, C. D.: Impact of brown and clear carbon on light absorption enhancement,  
779 single scatter albedo and absorption wavelength dependence of black carbon, *Atmos. Chem. Phys.*, 10,  
780 4207-4220, doi: DOI 10.5194/acp-10-4207-2010, 2010.

781 Lack, D. A., Langridge, J. M., Bahreini, R., Cappa, C. D., Middlebrook, A. M., and Schwarz, J. P.:  
782 Brown carbon and internal mixing in biomass burning particles, *P Natl Acad Sci USA*, 109, 14802-  
783 14807, doi: 10.1073/pnas.1206575109, 2012a.

784 Lack, D. A., Richardson, M. S., Law, D., Langridge, J. M., Cappa, C. D., McLaughlin, R. J., and  
785 Murphy, D. M.: Aircraft instrument for comprehensive characterization of aerosol optical properties,  
786 Part 2: black and brown carbon absorption and absorption enhancement measured with photo acoustic  
787 spectroscopy, *Aerosol. Sci. Technol.*, 46, 555-568, 2012b.

788 Lack, D. A. and Langridge, J. M.: On the attribution of black and brown carbon light absorption using  
789 the Ångström exponent, *Atmos. Chem. Phys.*, 13, 10535-10543, doi: 10.5194/acp-13-10535-2013,  
790 2013.

791 Lan, Z.-J., Huang, X.-F., Yu, K.-Y., Sun, T.-L., Zeng, L.-W., and Hu, M.: Light absorption of black  
792 carbon aerosol and its enhancement by mixing state in an urban atmosphere in South China, *Atmos.*  
793 *Environ.*, 69, 118-123, doi: 10.1016/j.atmosenv.2012.12.009, 2013.

794 Langridge, J. M., Richardson, M. S., Lack, D. A., Brock, C. A., and Murphy, D. M.: Limitations of  
795 the Photoacoustic Technique for Aerosol Absorption Measurement at High Relative Humidity,  
796 *Aerosol. Sci. Technol.*, 47, 1163-1173, doi: 10.1080/02786826.2013.827324, 2013.

797 Leung, K. K., Schnitzler, E. G., Jäger, W., and Olfert, J. S.: Relative Humidity Dependence of Soot  
798 Aggregate Restructuring Induced by Secondary Organic Aerosol: Effects of Water on Coating  
799 Viscosity and Surface Tension, *Environmental Science & Technology Letters*, doi:  
800 10.1021/acs.estlett.7b00298, 2017.

801 Lewis, K. A., Arnott, W. P., Moosmuller, H., Chakrabarty, R. K., Carrico, C. M., Kreidenweis, S. M.,  
802 Day, D. E., Malm, W. C., Laskin, A., Jimenez, J. L., Ulbrich, I. M., Huffman, J. A., Onasch, T. B.,  
803 Trimborn, A., Liu, L., and Mishchenko, M. I.: Reduction in biomass burning aerosol light absorption  
804 upon humidification: roles of inorganically-induced hygroscopicity, particle collapse, and  
805 photoacoustic heat and mass transfer, *Atmos. Chem. Phys.*, 9, 8949-8966, 2009a.

806 Lewis, K. A., Arnott, W. P., Moosmüller, H., Chakrabarty, R. K., Carrico, C. M., Kreidenweis, S. M.,  
807 Day, D. E., Malm, W. C., Laskin, A., Jimenez, J. L., Ulbrich, I. M., Huffman, J. A., Onasch, T. B.,  
808 Trimborn, A., Liu, L., and Mishchenko, M. I.: Reduction in biomass burning aerosol light absorption  
809 upon humidification: roles of inorganically-induced hygroscopicity, particle collapse, and  
810 photoacoustic heat and mass transfer, *Atmos. Chem. Phys.*, 9, 8949-8966, doi: 10.5194/acp-9-8949-  
811 2009, 2009b.

812 Lin, P., Aiona, P. K., Li, Y., Shiraiwa, M., Laskin, J., Nizkorodov, S. A., and Laskin, A.: Molecular  
813 Characterization of Brown Carbon in Biomass Burning Aerosol Particles, *Environ. Sci. Technol.*, 50,  
814 11815-11824, doi: 10.1021/acs.est.6b03024, 2016.

815 Liu, D., Allan, J., Whitehead, J., Young, D., Flynn, M., Coe, H., McFiggans, G., Fleming, Z. L., and  
816 Bandy, B.: Ambient black carbon particle hygroscopic properties controlled by mixing state and  
817 composition, *Atmos. Chem. Phys.*, 13, 2015-2029, doi: 10.5194/acp-13-2015-2013, 2013.

818 Liu, D., Whitehead, J., Alfarra, M. R., Reyes-Villegas, E., Spracklen, D. V., Reddington, C. L., Kong,  
819 S., Williams, P. I., Ting, Y.-C., Haslett, S., Taylor, J. W., Flynn, M. J., Morgan, W. T., McFiggans, G.,  
820 Coe, H., and Allan, J. D.: Black-carbon absorption enhancement in the atmosphere determined by  
821 particle mixing state, *Nature Geosci*, 10, 184-188, doi: 10.1038/ngeo2901, 2017.

822 Liu, F., Yon, J., and Bescond, A.: On the radiative properties of soot aggregates – Part 2: Effects of  
823 coating, *Journal of Quantitative Spectroscopy and Radiative Transfer*, 172, 134-145, doi:  
824 10.1016/j.jqsrt.2015.08.005, 2016a.

825 Liu, J., Lin, P., Laskin, A., Laskin, J., Kathmann, S. M., Wise, M., Caylor, R., Imholt, F., Selimovic,  
826 V., and Shilling, J. E.: Optical properties and aging of light-absorbing secondary organic aerosol,  
827 *Atmos. Chem. Phys.*, 16, 12815-12827, doi: 10.5194/acp-16-12815-2016, 2016b.

828 Liu, S., Aiken, A. C., Gorkowski, K., Dubey, M. K., Cappa, C. D., Williams, L. R., Herndon, S. C.,  
829 Massoli, P., Fortner, E. C., Chhabra, P. S., Brooks, W. A., Onasch, T. B., Jayne, J. T., Worsnop, D. R.,  
830 China, S., Sharma, N., Mazzoleni, C., Xu, L., Ng, N. L., Liu, D., Allan, J. D., Lee, J. D., Fleming, Z.  
831 L., Mohr, C., Zotter, P., Szidat, S., and Prevot, A. S. H.: Enhanced light absorption by mixed source  
832 black and brown carbon particles in UK winter, *Nat Commun*, 6, doi: 10.1038/ncomms9435, 2015.

833 Ma, N., Zhao, C. S., Muller, T., Cheng, Y. F., Liu, P. F., Deng, Z. Z., Xu, W. Y., Ran, L., Nekat, B.,  
834 van Pinxteren, D., Gnauk, T., Mueller, K., Herrmann, H., Yan, P., Zhou, X. J., and Wiedensohler, A.:  
835 A new method to determine the mixing state of light absorbing carbonaceous using the measured  
836 aerosol optical properties and number size distributions, *Atmos. Chem. Phys.*, 12, 2381-2397, doi:  
837 DOI 10.5194/acp-12-2381-2012, 2012.

838 Replacement Filter Tape for the Magee Scientific Model AE33 Aethalometer®:  
839 [http://www.mageesci.com/images/stories/docs/Magee\\_Scientific\\_Filter\\_Aethalometer\\_AE\\_Tape\\_Replacement\\_discussion.pdf](http://www.mageesci.com/images/stories/docs/Magee_Scientific_Filter_Aethalometer_AE_Tape_Replacement_discussion.pdf), 2017.

841 Matsui, H., Koike, M., Kondo, Y., Moteki, N., Fast, J. D., and Zaveri, R. A.: Development and  
842 validation of a black carbon mixing state resolved three-dimensional model: Aging processes and  
843 radiative impact, *J. Geophys. Res.*, 118, 2304-2326, doi: 10.1029/2012JD018446, 2013.

844 McMeeking, G. R., Good, N., Petters, M. D., McFiggans, G., and Coe, H.: Influences on the fraction  
845 of hydrophobic and hydrophilic black carbon in the atmosphere, *Atmos. Chem. Phys.*, 11, 5099-5112,  
846 doi: 10.5194/acp-11-5099-2011, 2011.

847 McMeeking, G. R., Fortner, E., Onasch, T. B., Taylor, J. W., Flynn, M., Coe, H., and Kreidenweis, S.  
848 M.: Impacts of nonrefractory material on light absorption by aerosols emitted from biomass burning,  
849 *J. Geophys. Res.*, 119, 12,272-212,286, doi: 10.1002/2014JD021750, 2014.

850 Moffet, R. C., O'Brien, R. E., Alpert, P. A., Kelly, S. T., Pham, D. Q., Gilles, M. K., Knopf, D. A., and  
851 Laskin, A.: Morphology and mixing of black carbon particles collected in central California during the  
852 CARES field study, *Atmos. Chem. Phys.*, 16, 14515-14525, doi: 10.5194/acp-16-14515-2016, 2016.

853 Moosmuller, H., Chakrabarty, R. K., Ehlers, K. M., and Arnott, W. P.: Absorption Angstrom  
854 coefficient, brown carbon, and aerosols: basic concepts, bulk matter, and spherical particles, *Atmos.*  
855 *Chem. Phys.*, 11, 1217-1225, doi: DOI 10.5194/acp-11-1217-2011, 2011.

856 Moteki, N., Kondo, Y., and Adachi, K.: Identification by single-particle soot photometer of black  
857 carbon particles attached to other particles: Laboratory experiments and ground observations in Tokyo,  
858 *J. Geophys. Res.*, 119, 2013JD020655, doi: 10.1002/2013jd020655, 2014.

859 Nakayama, T., Ikeda, Y., Sawada, Y., Setoguchi, Y., Ogawa, S., Kawana, K., Mochida, M., Ikemori,  
860 F., Matsumoto, K., and Matsumi, Y.: Properties of light-absorbing aerosols in the Nagoya urban area,  
861 Japan, in August 2011 and January 2012: Contributions of brown carbon and lensing effect, *J. Geophys.*  
862 *Res.*, 119, 2014JD021744, doi: 10.1002/2014JD021744, 2014.

863 Naoe, H., Hasegawa, S., Heintzenberg, J., Okada, K., Uchiyama, A., Zaizen, Y., Kobayashi, E., and  
864 Yamazaki, A.: State of mixture of atmospheric submicrometer black carbon particles and its effect on  
865 particulate light absorption, *Atmos. Environ.*, 43, 1296-1301, doi: 10.1016/j.atmosenv.2008.11.031,  
866 2009.

867 Nessler, R., Weingartner, E., and Baltensperger, U.: Effect of humidity on aerosol light absorption and  
868 its implications for extinction and the single scattering albedo illustrated for a site in the lower free  
869 troposphere, *J. Aerosol. Sci.*, 36, 958-972, doi: 10.1016/j.jaerosci.2004.11.012, 2005.

870 Nordmann, S., Cheng, Y. F., Carmichael, G. R., Yu, M., Denier van der Gon, H. A. C., Zhang, Q.,  
871 Saide, P. E., Pöschl, U., Su, H., Birmili, W., and Wiedensohler, A.: Atmospheric black carbon and  
872 warming effects influenced by the source and absorption enhancement in central Europe, *Atmos. Chem.*  
873 *Phys.*, 14, 12683-12699, doi: 10.5194/acp-14-12683-2014, 2014.

874 Pandey, A., Pervez, S., and Chakrabarty, R. K.: Filter-based measurements of UV-vis mass absorption  
875 cross sections of organic carbon aerosol from residential biomass combustion: Preliminary findings  
876 and sources of uncertainty, *Journal of Quantitative Spectroscopy and Radiative Transfer*, 182, 296-  
877 304, doi: 10.1016/j.jqsrt.2016.06.023, 2016.

878 Peng, J., Hu, M., Guo, S., Du, Z., Zheng, J., Shang, D., Levy Zamora, M., Zeng, L., Shao, M., Wu, Y.-  
879 S., Zheng, J., Wang, Y., Glen, C. R., Collins, D. R., Molina, M. J., and Zhang, R.: Markedly enhanced

880 absorption and direct radiative forcing of black carbon under polluted urban environments,  
881 Proceedings of the National Academy of Sciences, 113, 4266-4271, doi: 10.1073/pnas.1602310113,  
882 2016.

883 Pokhrel, R. P., Beamesderfer, E. R., Wagner, N. L., Langridge, J. M., Lack, D. A., Jayarathne, T.,  
884 Stone, E. A., Stockwell, C. E., Yokelson, R. J., and Murphy, S. M.: Relative importance of black  
885 carbon, brown carbon, and absorption enhancement from clear coatings in biomass burning emissions,  
886 Atmos. Chem. Phys., 17, 5063-5078, doi: 10.5194/acp-17-5063-2017, 2017.

887 Ramanathan, V. and Carmichael, G.: Global and regional climate changes due to black carbon, Nat  
888 Geosci, 1, 221-227, doi: Doi 10.1038/Ngeo156, 2008.

889 Raspert, R., Slaton, W. V., Arnott, W. P., and Moosmüller, H.: Evaporation–Condensation Effects on  
890 Resonant Photoacoustics of Volatile Aerosols, Journal of Atmospheric and Oceanic Technology, 20,  
891 685-695, doi: 10.1175/1520-0426(2003)20<685:ecorp>2.0.co;2, 2003.

892 Reid, J. S., Eck, T. F., Christopher, S. A., Koppmann, R., Dubovik, O., Eleuterio, D. P., Holben, B. N.,  
893 Reid, E. A., and Zhang, J.: A review of biomass burning emissions part III: intensive optical properties  
894 of biomass burning particles, Atmos. Chem. Phys., 5, 827-849, doi: 10.5194/acp-5-827-2005, 2005.

895 Roden, C. A., Bond, T. C., Conway, S., and Pinel, A. B. O.: Emission factors and real-time optical  
896 properties of particles emitted from traditional wood burning cookstoves, Environ. Sci. Technol., 40,  
897 6750-6757, doi: 10.1021/es052080i, 2006.

898 Rose, D., Wehner, B., Ketzler, M., Engler, C., Voigtländer, J., Tuch, T., and Wiedensohler, A.:  
899 Atmospheric number size distributions of soot particles and estimation of emission factors, Atmos.  
900 Chem. Phys., 6, 1021-1031, doi: 10.5194/acp-6-1021-2006, 2006.

901 Saathoff, H., Naumann, K. H., Schnaiter, M., Schöck, W., Möhler, O., Schurath, U., Weingartner, E.,  
902 Gysel, M., and Baltensperger, U.: Coating of soot and (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub> particles by ozonolysis products of  
903  $\alpha$ -pinene, J. Aerosol. Sci., 34, 1297-1321, doi: 10.1016/S0021-8502(03)00364-1, 2003.

904 Sandradewi, J., Prévôt, A. S. H., Weingartner, E., Schmidhauser, R., Gysel, M., and Baltensperger, U.:  
905 A study of wood burning and traffic aerosols in an Alpine valley using a multi-wavelength  
906 Aethalometer, Atmos. Environ., 42, 101-112, doi: 10.1016/j.atmosenv.2007.09.034, 2008.

907 Saturno, J., Pöhlker, C., Massabò, D., Brito, J., Carbone, S., Cheng, Y., Chi, X., Ditas, F., Hrabě de  
908 Angelis, I., Morán-Zuloaga, D., Pöhlker, M. L., Rizzo, L. V., Walter, D., Wang, Q., Artaxo, P., Prati,  
909 P., and Andreae, M. O.: Comparison of different Aethalometer correction schemes and a reference  
910 multi-wavelength absorption technique for ambient aerosol data, Atmos. Meas. Tech., 10, 2837-2850,  
911 doi: 10.5194/amt-10-2837-2017, 2017.

912 Schmid, O., Artaxo, P., Arnott, W. P., Chand, D., Gatti, L. V., Frank, G. P., Hoffer, A., Schnaiter, M.,  
913 and Andreae, M. O.: Spectral light absorption by ambient aerosols influenced by biomass burning in  
914 the Amazon Basin. I: Comparison and field calibration of absorption measurement techniques, Atmos.  
915 Chem. Phys., 6, 3443-3462, 2006.

916 Schnaiter, M., Linke, C., Mohler, O., Naumann, K. H., Saathoff, H., Wagner, R., Schurath, U., and  
917 Wehner, B.: Absorption amplification of black carbon internally mixed with secondary organic aerosol,  
918 J. Geophys. Res., 110, -, 2005.

919 Schwarz, J. P., Spackman, J. R., Fahey, D. W., Gao, R. S., Lohmann, U., Stier, P., Watts, L. A.,  
920 Thomson, D. S., Lack, D. A., Pfister, L., Mahoney, M. J., Baumgardner, D., Wilson, J. C., and Reeves,

921 J. M.: Coatings and their enhancement of black carbon light absorption in the tropical atmosphere, *J.*  
922 *Geophys. Res.*, 113, -, 2008.

923 Sedlacek, A. J., Lewis, E. R., Kleinman, L., Xu, J. Z., and Zhang, Q.: Determination of and evidence  
924 for non-core-shell structure of particles containing black carbon using the Single-Particle Soot  
925 Photometer (SP2), *Geophys. Res. Lett.*, 39, 2012.

926 Shen, G., Chen, Y., Wei, S., Fu, X., Zhu, Y., and Tao, S.: Mass absorption efficiency of elemental  
927 carbon for source samples from residential biomass and coal combustions, *Atmos. Environ.*, 79, 79-  
928 84, doi: 10.1016/j.atmosenv.2013.05.082, 2013.

929 Shiraiwa, M., Kondo, Y., Iwamoto, T., and Kita, K.: Amplification of Light Absorption of Black  
930 Carbon by Organic Coating, *Aerosol. Sci. Technol.*, 44, 46-54, 2010.

931 Suglia, S. F., Gryparis, A., Wright, R. O., Schwartz, J., and Wright, R. J.: Association of Black Carbon  
932 with Cognition among Children in a Prospective Birth Cohort Study, *American Journal of*  
933 *Epidemiology*, 167, 280-286, doi: 10.1093/aje/kwm308, 2008.

934 Tan, H., Liu, L., Fan, S., Li, F., Yin, Y., Cai, M., and Chan, P. W.: Aerosol optical properties and  
935 mixing state of black carbon in the Pearl River Delta, China, *Atmos. Environ.*, 131, 196-208, doi:  
936 10.1016/j.atmosenv.2016.02.003, 2016.

937 Tao, W. K., Chen, J. P., Li, Z. Q., Wang, C., and Zhang, C. D.: Impact of Aerosols on Convective  
938 Clouds and Precipitation, *Rev Geophys*, 50, Rg2001, doi: Doi 10.1029/2011rg000369, 2012.

939 Tavakoli, F. and Olfert, J. S.: Determination of particle mass, effective density, mass–mobility  
940 exponent, and dynamic shape factor using an aerodynamic aerosol classifier and a differential mobility  
941 analyzer in tandem, *J. Aerosol. Sci.*, 75, 35-42, doi: 10.1016/j.jaerosci.2014.04.010, 2014.

942 ten Brink, H., Otjes, R., Jongejan, P., and Slanina, S.: An instrument for semi-continuous monitoring  
943 of the size-distribution of nitrate, ammonium, sulphate and chloride in aerosol, *Atmos. Environ.*, 41,  
944 2768-2779, doi: 10.1016/j.atmosenv.2006.11.041, 2007.

945 Turpin, B. J. and Huntzicker, J. J.: Secondary Formation of Organic Aerosol in the Los-Angeles Basin  
946 - a Descriptive Analysis of Organic and Elemental Carbon Concentrations, *Atmos. Environ.*, 25, 207-  
947 215, 1991.

948 Ueda, S., Nakayama, T., Taketani, F., Adachi, K., Matsuki, A., Iwamoto, Y., Sadanaga, Y., and  
949 Matsumi, Y.: Light absorption and morphological properties of soot-containing aerosols observed at  
950 an East Asian outflow site, Noto Peninsula, Japan, *Atmos. Chem. Phys.*, 16, 2525-2541, doi:  
951 10.5194/acp-16-2525-2016, 2016.

952 Virkkula, A., Makela, T., Hillamo, R., Yli-Tuomi, T., Hirsikko, A., Hameri, K., and Koponen, I. K.:  
953 A simple procedure for correcting loading effects of aethalometer data, *J. Air Waste Manage. Assoc.*,  
954 57, 1214-1222, 2007.

955 Wang, Q., Huang, R., Zhao, Z., Cao, J., Ni, H., Tie, X., Zhu, C., Shen, Z., Wang, M., and Dai, W.:  
956 Effects of photochemical oxidation on the mixing state and light absorption of black carbon in the  
957 urban atmosphere of China, *Environmental Research Letters*, 12, 044012, 2017.

958 Wang, Q. Y., Huang, R. J., Cao, J. J., Han, Y. M., Wang, G. H., Li, G. H., Wang, Y. C., Dai, W. T.,  
959 Zhang, R. J., and Zhou, Y. Q.: Mixing State of Black Carbon Aerosol in a Heavily Polluted Urban  
960 Area of China: Implications for Light Absorption Enhancement, *Aerosol. Sci. Technol.*, 48, 689-697,  
961 doi: 10.1080/02786826.2014.917758, 2014.

962 Wang, Y. Q.: *MeteoInfo: GIS software for meteorological data visualization and analysis*,  
963 *Meteorological Applications*, 21, 360-368, doi: 10.1002/met.1345, 2014.

964 Weingartner, E., Saathoff, H., Schnaiter, M., Streit, N., Bitnar, B., and Baltensperger, U.: Absorption  
965 of light by soot particles: determination of the absorption coefficient by means of aethalometers, *J.*  
966 *Aerosol. Sci.*, 34, 1445-1463, doi: 10.1016/S0021-8502(03)00359-8, 2003.

967 Weyant, C. L., Shepson, P. B., Subramanian, R., Cambaliza, M. O. L., Heimbürger, A., McCabe, D.,  
968 Baum, E., Stirm, B. H., and Bond, T. C.: Black Carbon Emissions from Associated Natural Gas Flaring,  
969 *Environ. Sci. Technol.*, 50, 2075-2081, doi: 10.1021/acs.est.5b04712, 2016.

970 Wild, M.: Enlightening Global Dimming and Brightening, *B Am Meteorol Soc*, 93, 27-37, doi:  
971 10.1175/bams-d-11-00074.1, 2011.

972 Wu, C., Ng, W. M., Huang, J., Wu, D., and Yu, J. Z.: Determination of Elemental and Organic Carbon  
973 in PM<sub>2.5</sub> in the Pearl River Delta Region: Inter-Instrument (Sunset vs. DRI Model 2001  
974 Thermal/Optical Carbon Analyzer) and Inter-Protocol Comparisons (IMPROVE vs. ACE-Asia  
975 Protocol), *Aerosol. Sci. Technol.*, 46, 610-621, doi: 10.1080/02786826.2011.649313, 2012.

976 Wu, C., Huang, X. H. H., Ng, W. M., Griffith, S. M., and Yu, J. Z.: Inter-comparison of NIOSH and  
977 IMPROVE protocols for OC and EC determination: implications for inter-protocol data conversion,  
978 *Atmos. Meas. Tech.*, 9, 4547-4560, doi: 10.5194/amt-9-4547-2016, 2016a.

979 Wu, C. and Yu, J. Z.: Determination of primary combustion source organic carbon-to-elemental carbon  
980 (OC/EC) ratio using ambient OC and EC measurements: secondary OC-EC correlation minimization  
981 method, *Atmos. Chem. Phys.*, 16, 5453-5465, doi: 10.5194/acp-16-5453-2016, 2016.

982 Wu, D., Mao, J. T., Deng, X. J., Tie, X. X., Zhang, Y. H., Zeng, L. M., Li, F., Tan, H. B., Bi, X. Y.,  
983 Huang, X. Y., Chen, J., and Deng, T.: Black carbon aerosols and their radiative properties in the Pearl  
984 River Delta region, *Sci China Ser D*, 52, 1152-1163, doi: 10.1007/s11430-009-0115-y, 2009.

985 Wu, D., Wu, C., Liao, B., Chen, H., Wu, M., Li, F., Tan, H., Deng, T., Li, H., Jiang, D., and Yu, J. Z.:  
986 Black carbon over the South China Sea and in various continental locations in South China, *Atmos.*  
987 *Chem. Phys.*, 13, 12257-12270, doi: 10.5194/acp-13-12257-2013, 2013.

988 Wu, Y., Zhang, R., Tian, P., Tao, J., Hsu, S. C., Yan, P., Wang, Q., Cao, J., Zhang, X., and Xia, X.:  
989 Effect of ambient humidity on the light absorption amplification of black carbon in Beijing during  
990 January 2013, *Atmos. Environ.*, 124, Part B, 217-223, doi: 10.1016/j.atmosenv.2015.04.041, 2016b.

991 Yang, M., Howell, S. G., Zhuang, J., and Huebert, B. J.: Attribution of aerosol light absorption to black  
992 carbon, brown carbon, and dust in China - interpretations of atmospheric measurements during EAST-  
993 AIRE, *Atmos. Chem. Phys.*, 9, 2035-2050, 2009.

994 Yu, H., Wu, C., Wu, D., and Yu, J. Z.: Size distributions of elemental carbon and its contribution to  
995 light extinction in urban and rural locations in the pearl river delta region, China, *Atmos. Chem. Phys.*,  
996 10, 5107-5119, doi: 10.5194/acp-10-5107-2010, 2010.

997 Yuan, J. F., Huang, X. F., Cao, L. M., Cui, J., Zhu, Q., Huang, C. N., Lan, Z. J., and He, L. Y.: Light  
998 absorption of brown carbon aerosol in the PRD region of China, *Atmos. Chem. Phys.*, 16, 1433-1443,  
999 doi: 10.5194/acp-16-1433-2016, 2016.

1000 Zhang, G., Bi, X., Qiu, N., Han, B., Lin, Q., Peng, L., Chen, D., Wang, X., Peng, P., Sheng, G., and  
1001 Zhou, Z.: The real part of the refractive indices and effective densities for chemically segregated  
1002 ambient aerosols in Guangzhou measured by a single-particle aerosol mass spectrometer, *Atmos.*  
1003 *Chem. Phys.*, 16, 2631-2640, doi: 10.5194/acp-16-2631-2016, 2016a.



1004 Zhang, R. Y., Khalizov, A. F., Pagels, J., Zhang, D., Xue, H. X., and McMurry, P. H.: Variability in  
1005 morphology, hygroscopicity, and optical properties of soot aerosols during atmospheric processing, *P*  
1006 *Natl Acad Sci USA*, 105, 10291-10296, 2008.

1007 Zhang, Y., Zhang, Q., Cheng, Y., Su, H., Kecorius, S., Wang, Z., Wu, Z., Hu, M., Zhu, T.,  
1008 Wiedensohler, A., and He, K.: Measuring the morphology and density of internally mixed black carbon  
1009 with SP2 and VTDMA: new insight into the absorption enhancement of black carbon in the atmosphere,  
1010 *Atmos. Meas. Tech.*, 9, 1833-1843, doi: 10.5194/amt-9-1833-2016, 2016b.

1011

1012

1013 Table 1. Abbreviations.

1014

Abbreviation	Definition
AAE <sub>470-660</sub>	Absorption Angstrom Exponent between 470 and 660 nm
BB	Biomass burning
BrC	Brown Carbon
D <sub>core</sub> , D <sub>shell</sub>	Particle diameter of core/shell
E <sub>abs550</sub>	Light absorption enhancement factor at 550 nm
$\sigma_{abs550}$	Light absorption coefficient at 550 nm
$\sigma_{abs,t}$	Total light absorption coefficient of a coated particle
$\sigma_{abs,p}$	Primary light absorption coefficient attributed to the soot core alone of a coated particle
$\sigma_{abs,c}$	Extra light absorption coefficient due to the lensing effect of coating on the soot core
LII	Laser induced incandescence technique for soot measurement
LWC	Liquid water content
MAE <sub>550</sub>	Mass absorption efficiency at 550 nm, also known as mass absorption cross-section (MAC)
MAE <sub>p,550</sub>	Primary MAE of freshly emitted soot particles at 550 nm
MAAP	Multi Angle Absorption Photometer
MOUDI	Micro Orifice Uniform Deposit Impactor
MRS	Minimum R squared method
PAS	Photo acoustic spectrometer
PRD	Pearl River Delta region, China
SP2	Single particle soot photometer
SSA	Single scattering albedo
TD	Thermal denuder
TOA	Thermal optical analysis
TSV	Total spatial variance in backward trajectories cluster analysis

1015

1016 Table 2. Comparison of MRS application on  $(OC/EC)_p$  (for SOC estimation) and  $MAE_p$  (for  $E_{abs}$  estimation).  
 1017

	MRS in EC tracer method for SOC estimation (Wu and Yu, 2016)	MRS in EC tracer method for $E_{abs}$ estimation (this study)
Key parameter of fresh EC particles to be determined	$\left(\frac{OC}{EC}\right)_p = \frac{POC}{EC}$	$MAE_p = \frac{O\sigma_{abs,p}}{EC}$
Input quantities for MRS from measurements	OC, EC (tracer)	$\sigma_{abs,t}$ , EC (tracer)
Variable to be decoupled by the tracer	$OC = POC + SOC$ $= \left(\frac{OC}{EC}\right)_p \times EC + SOC$	$\sigma_{abs,t} = \sigma_{abs,p} + \sigma_{abs,c}$ $= \left(\frac{O\sigma_{abs,t}C}{EC}\right)_p \times EC + \sigma_{abs,c}$
Ambient measurement at its closest to fresh emissions	Minimum $R^2$ (SOC, EC) $SOC = OC - \left(\frac{OC}{EC}\right)_p \times EC$	Minimum $R^2$ ( $\sigma_{abs,c}$ , EC) $\sigma_{abs,c} = \sigma_{abs,t} - MAE_p \times EC$
Graph		

1018

1019 Table 3. Comparison of  $E_{\text{abs}}$  between various studies.

1020

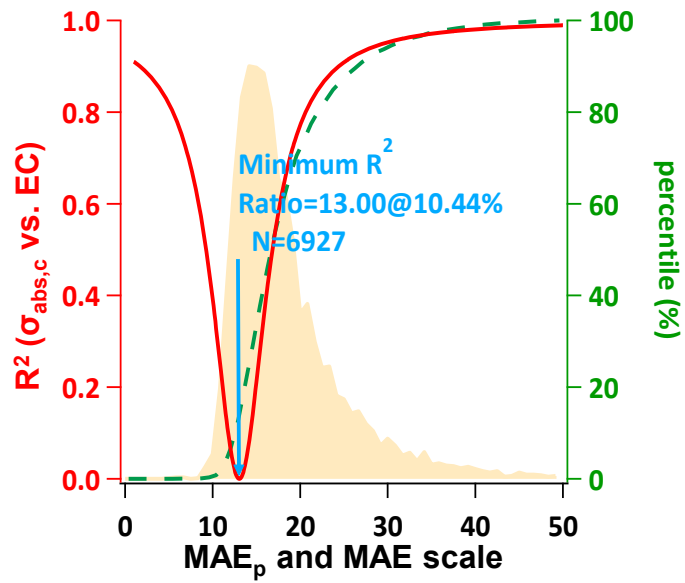
Location	Type	Sampling Duration	$\lambda$ (nm)	Instrument	$E_{\text{abs}}$	Method	Reference
Guangzhou, China	Suburban	2012.2-2013.1	550	AE+OCEC	1.50±0.48	MAE	This study
Xi'an, China	Urban	2012.12-2013.1	870	PAS	1.8	MAE	(Wang et al., 2014)
Shenzhen, China	Urban	2011.8-9	532	PAS	1.3	MAE	(Lan et al., 2013)
Jinan, China	Urban	2014.2	678	OCEC	2.07 ± 0.72	AFD	(Chen et al., 2017)
Nanjing, China	Suburban	2012.11	532	PAS	1.6	MAE	(Cui et al., 2016a)
Boulder, USA	Forest fire	2010.9	532	PAS	1.38	TD 200°C	(Lack et al., 2012a)
London, UK	Rural	2012.2	781	PAS	1.4	TD 250°C	(Liu et al., 2015)
California, USA	Rural	2010.6	532	PAS	1.06	TD 250°C	(Cappa et al., 2012)
Noto Peninsula, Japan	Rural	2013.4-5	781	PAS	1.22	TD 300°C	(Ueda et al., 2016)
Yuncheng, China	Rural	2014.6-7	678	OCEC	2.25 ± 0.55	AFD	(Cui et al., 2016b)
San Jose, Costa Rica	Rural	2006 winter	1064	SP2	1.3	Mie+SP2	(Schwarz et al., 2008)

1021

AE: Aethalometer ; OCEC: OCEC analyzer; PAS: photo acoustic spectrometer; SP2: Single particle soot photometer; TD: Thermal denuder AFD: filter filtration-dissolution

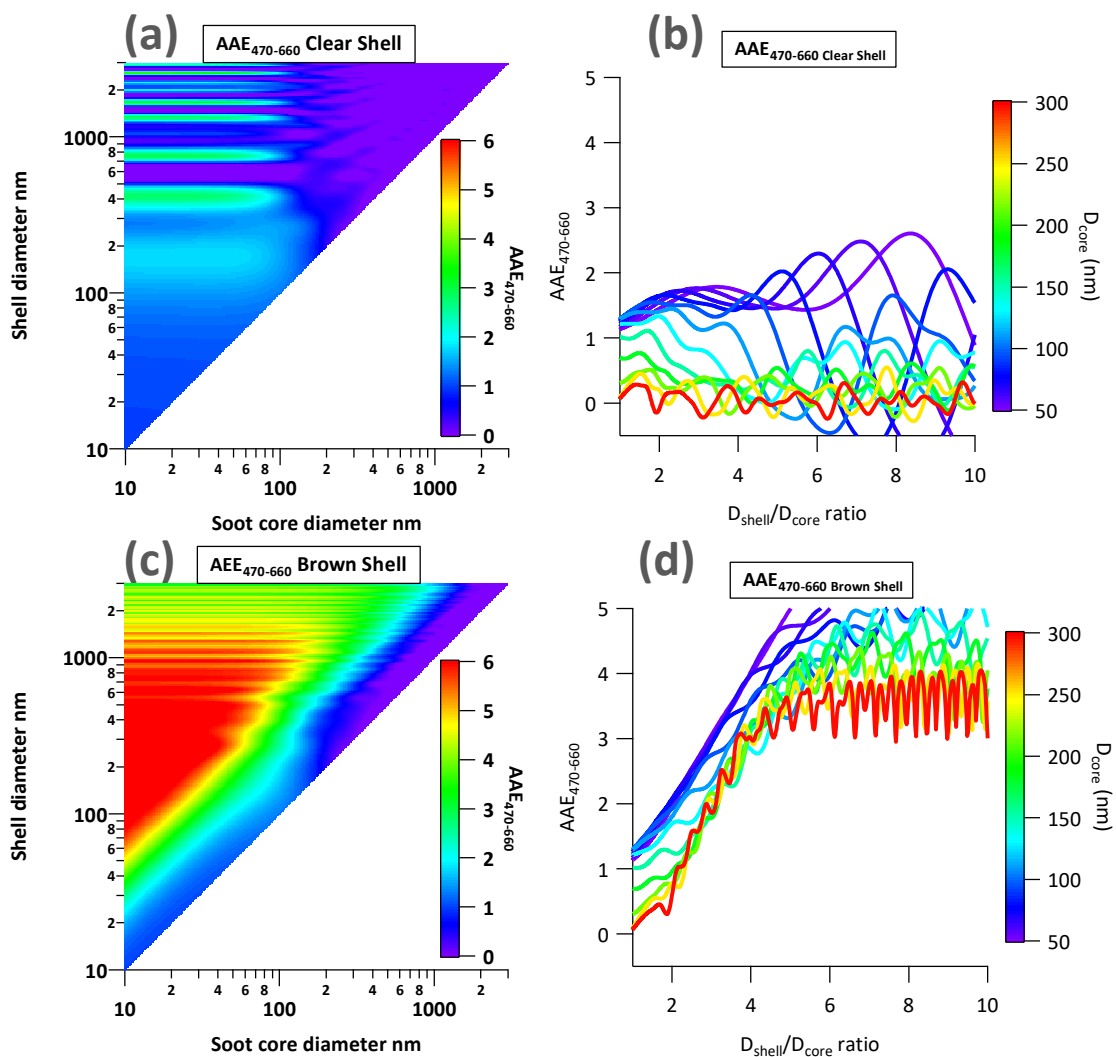
1022

1023



1024

1025 Figure 1. Minimum R squared (MRS) plot for calculating MAE<sub>p</sub> at 550 nm. The red curve is the correlation result between  
 1026  $\sigma_{abs,c}$  ( $\sigma_{abs,t} - EC * MAE_p$ ) and EC mass. The shaded area in light tan represents the frequency distribution of observed  
 1027 MAE. The dashed green line is the cumulative distribution of observed MAE.



1028

1029

1030

1031

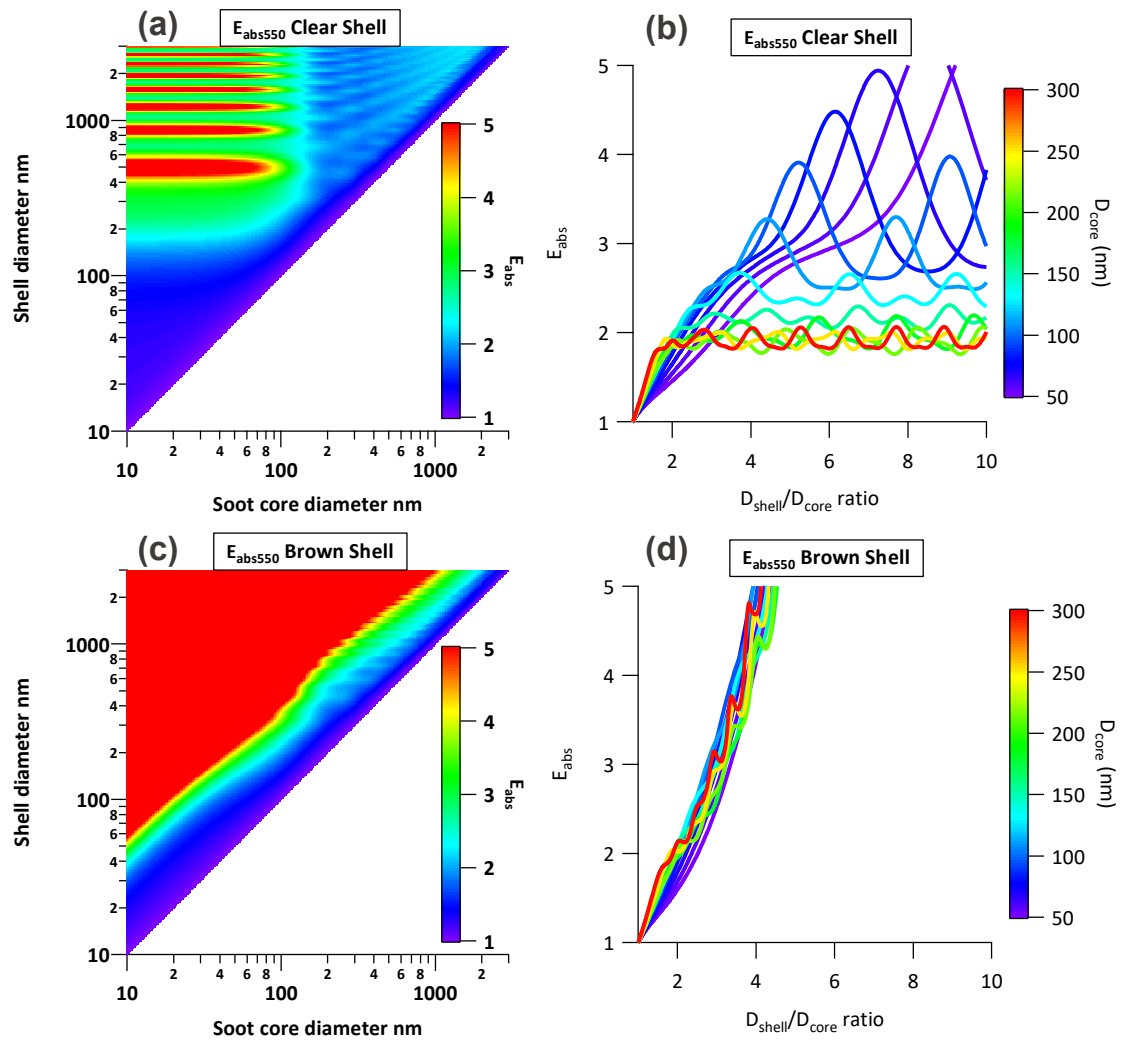
1032

1033

1034

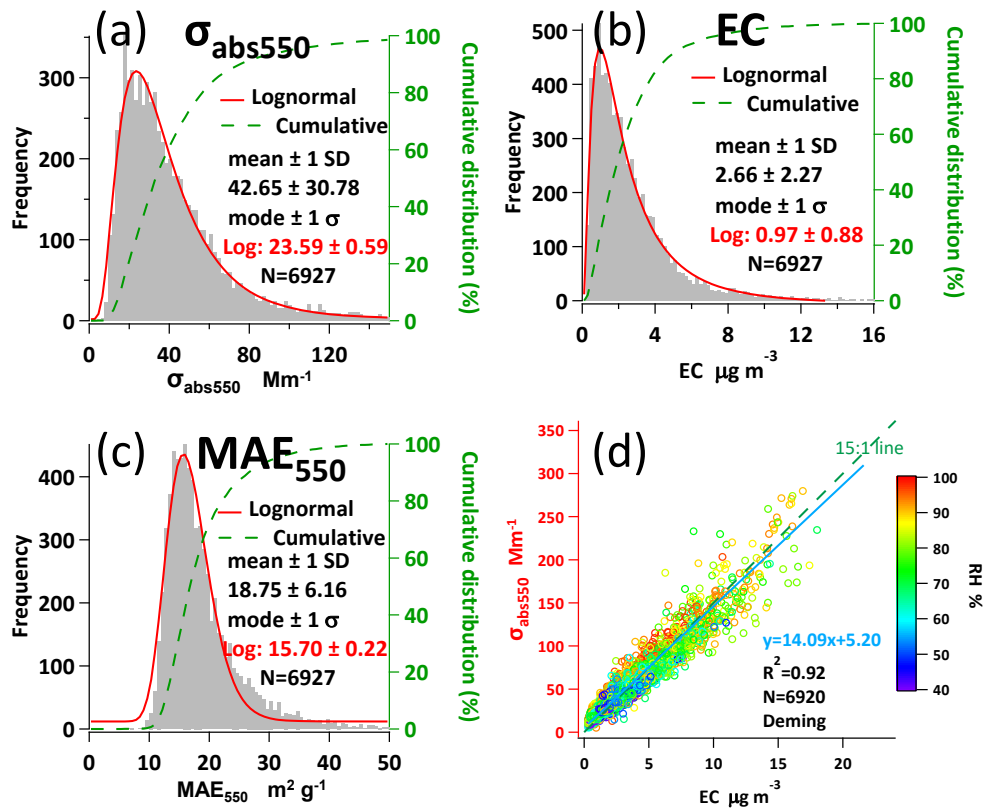
1035

Figure 2. Mie simulated size dependency of soot particles AAE<sub>470-660</sub>. (a) Combination of different clear shell (y axis) and core diameters (x axis). The color coding represents the AAE<sub>470-660</sub> of a particle with specific core and clear shell size; (b) Cross-sections views of (a). The color coding represents different D<sub>core</sub> in the range of 50 ~ 300 nm. (c)&(d) Similar to (a)&(b) but from the brown shell scenario.



1036  
 1037  
 1038  
 1039  
 1040  
 1041  
 1042  
 1043

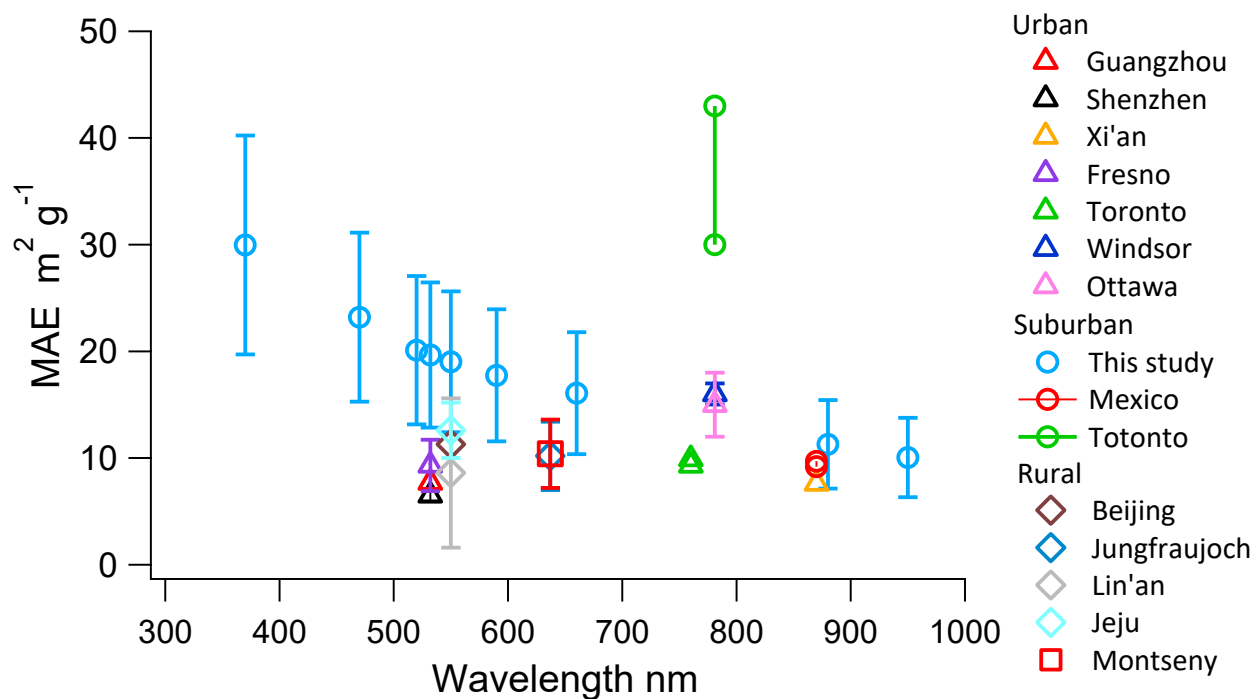
Figure 3. Mie simulated size dependency of soot particles  $E_{abs}$  at wavelength 550 nm. (a) Combination of different clear shell (y axis) and core diameters (x axis). The color coding represents the  $E_{abs}$  of a particle with specific core and clear shell size; (b) Cross-sections views of (a). The color coding represents different  $D_{core}$  in the range of 50 – 300 nm. (c)&(d) Similar to (a)&(b) but from the brown shell scenario.



1044

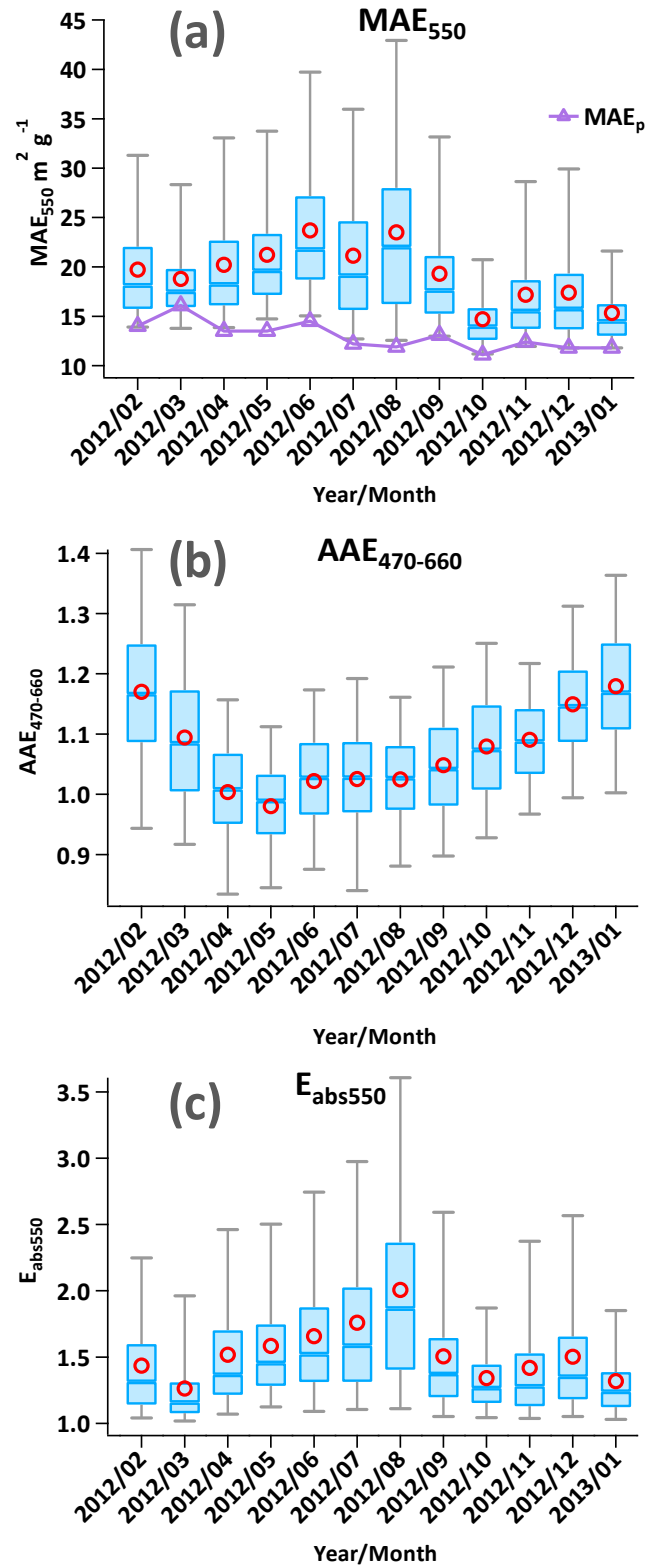
1045 Figure 4. Measured annual statistics of  $\sigma_{abs550}$ , EC and MAE<sub>550</sub>. (a) Annual frequency distribution of light absorption at  
 1046 550 nm. The red curve represents the fitting line for a log-normal distribution. (b) Annual frequency distribution of EC  
 1047 mass concentration (c) Frequency distribution of Mass absorption efficiency (MAE) at 550 nm. (d) Scatter plot of light  
 1048 absorption (550 nm) and EC mass. The slope represents MAE<sub>550</sub>. The blue regression line is by Deming regression. The  
 1049 color coding represents RH.  
 1050





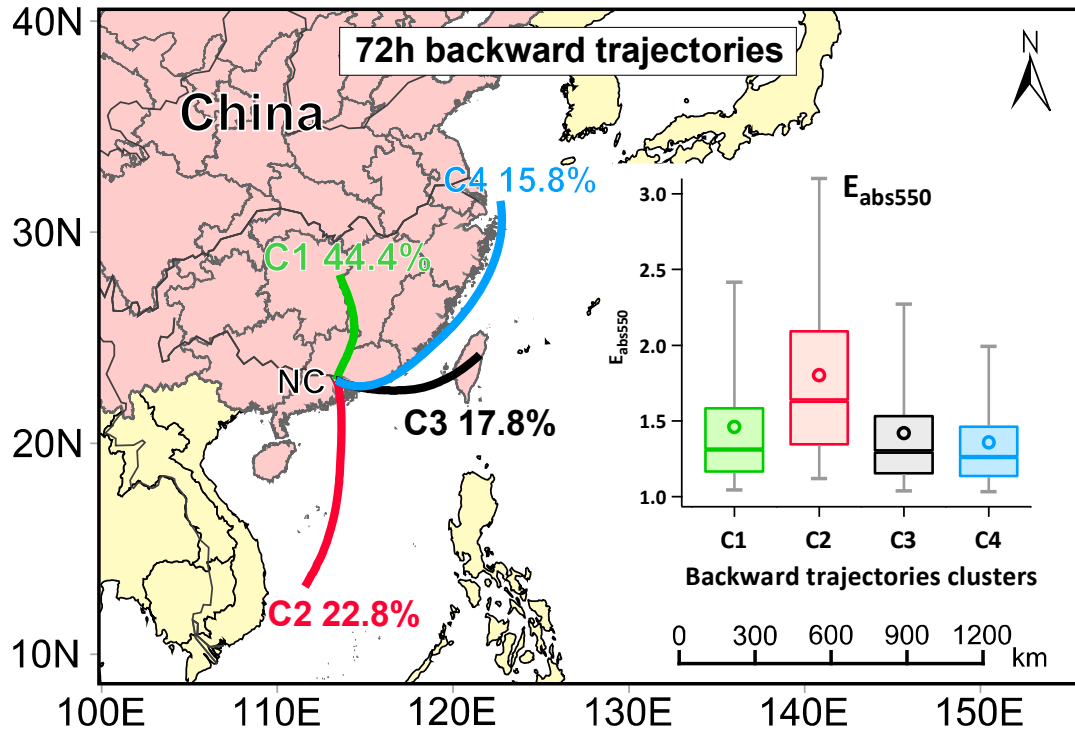
1051

1052 Figure 5. Comparison of spectral MAE measurements from this study with previous studies. Triangle, circle and rhombus  
 1053 represent urban, suburban and rural respectively. Details and reference can be found in Table S1. The whiskers represent  
 1054 one standard deviation.



1055

1056 Figure 6. Measured monthly variations of (a) MAE<sub>550</sub>, the purple line represents MAE<sub>p</sub> estimated by MRS (b) AAE<sub>470-660</sub>  
 1057 and (c) E<sub>abs550</sub>. Red circles represent the monthly average. The line inside the box indicates the monthly median. Upper and  
 1058 lower boundaries of the box represent the 75<sup>th</sup> and the 25<sup>th</sup> percentiles; the whiskers above and below each box represent  
 1059 the 95<sup>th</sup> and 5<sup>th</sup> percentiles.

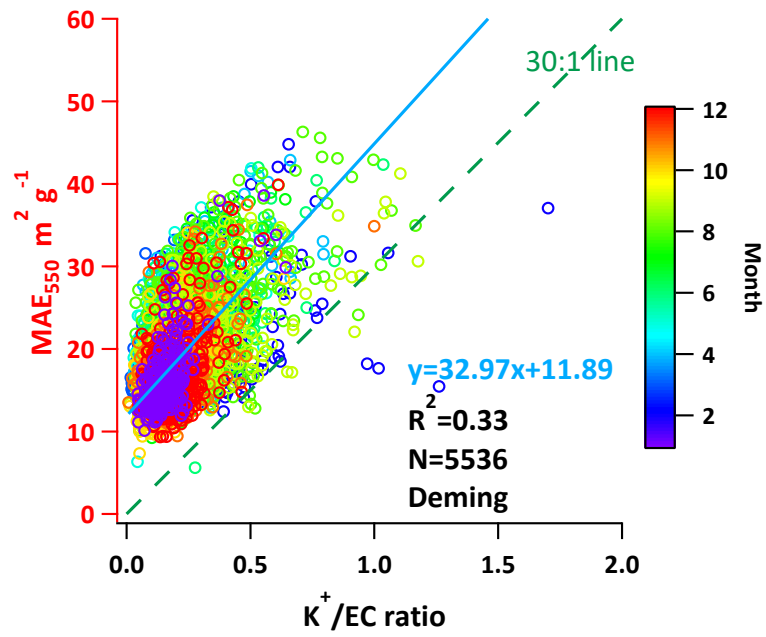


1060

1061

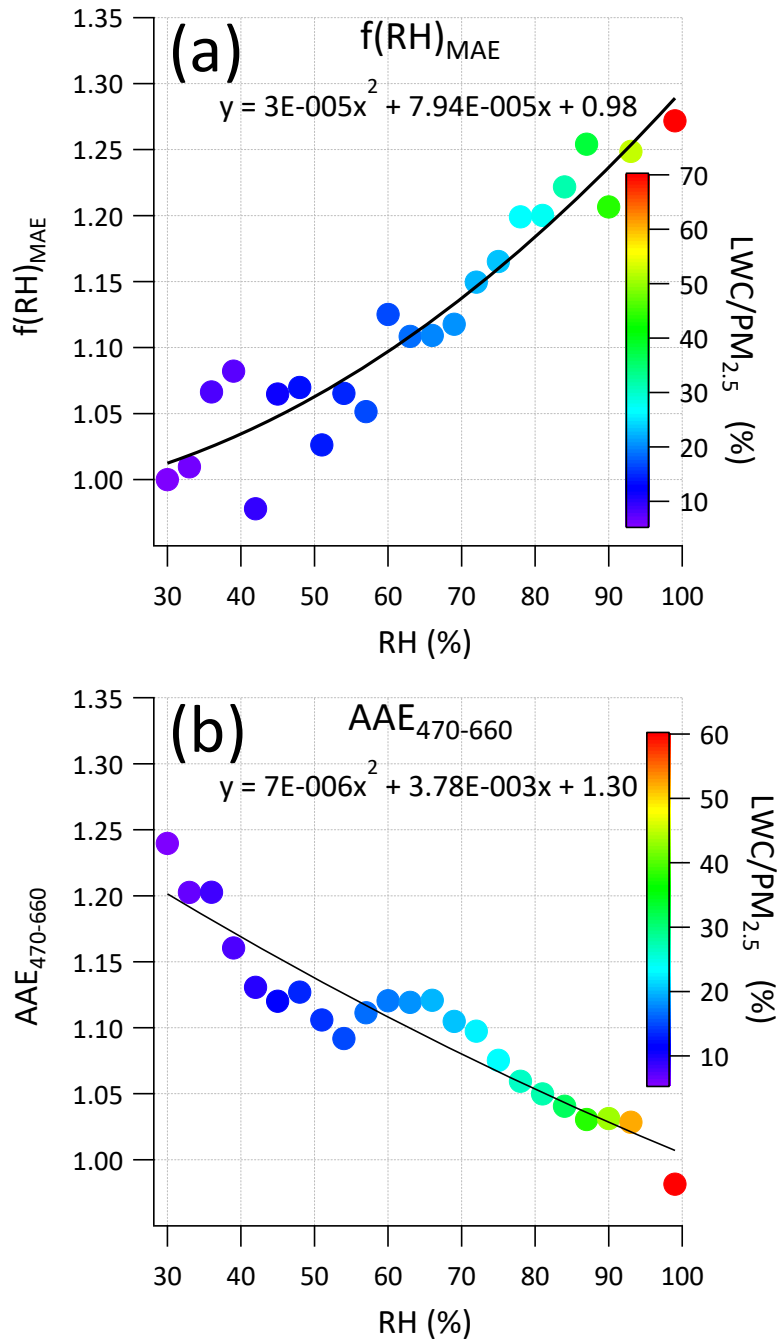
1062

Figure 7. Average backward trajectories arriving at 100 m at NC site for four clusters (2012 Feb - 2013 Jan).  $E_{abs550}$  by different clusters are shown in the box plot.



1063

1064 Figure 8. MAE<sub>550</sub> dependency on biomass burning indicator  $K^+/EC$  ratio. The color coding represents months. The intercept  
 1065 represents MAE without biomass burning effect. The 30:1 line serves as a reference line with an integer slope that  
 1066 is close to the regressed slope through the origin.



1067

1068

1069

Figure 9. Optical properties dependency on RH determined from one year's sampling data at NC site. (a) Hygroscopic growth factor ( $f(RH)$ ) of EC MAE<sub>550</sub> (b)  $AAE_{470-660}$  as a function of RH.

1 *Supplement of*

2 **Quantifying black carbon light absorption enhancement by**  
3 **a novel statistical approach**

4

5 **Cheng Wu<sup>1,2</sup>, Dui Wu<sup>1,2,3</sup>, Jian Zhen Yu<sup>4,5,6</sup>**

6

7 [1] Institute of Mass Spectrometer and Atmospheric Environment, Jinan University,  
8 Guangzhou 510632, China

9 [2] Guangdong Provincial Engineering Research Center for on-line source apportionment  
10 system of air pollution, Guangzhou 510632, China

11 [3] Institute of Tropical and Marine Meteorology, China Meteorological Administration,  
12 Guangzhou 510080, China

13 [4] Division of Environment, Hong Kong University of Science and Technology, Clear Water  
14 Bay, Hong Kong, China

15 [5] Atmospheric Research Centre, Fok Ying Tung Graduate School, Hong Kong University  
16 of Science and Technology, Nansha, China

17 [6] Department of Chemistry, Hong Kong University of Science and Technology, Clear Water  
18 Bay, Hong Kong, China

19 *Corresponding to:* Cheng Wu ([wucheng.vip@foxmail.com](mailto:wucheng.vip@foxmail.com)) and Jian Zhen Yu ([jian.yu@ust.hk](mailto:jian.yu@ust.hk))

20

21

22  
23 This SI contains five tables and twenty-five figures.

24

## 25 **Uncertainty of $E_{abs}$ estimation**

26 The uncertainty of  $E_{abs}$  estimation depends on uncertainty propagation from MAE uncertainty,  
27 which can be calculated from (Harris, 2010):

$$28 \quad MAE_{Unc} = MAE \times \sqrt{\left(\frac{\sigma_{abs,Unc}}{\sigma_{abs}}\right)^2 + \left(\frac{EC_{Unc}}{EC}\right)^2} \quad S1$$

$$29 \quad E_{abs,Unc} = E_{abs} \times \sqrt{\left(\frac{MAE_{Unc}}{MAE}\right)^2 + \left(\frac{MAE_{p,Unc}}{MAE_p}\right)^2} \quad S2$$

30

## 31 **Descriptions of customized programs used in this study for data analysis and** 32 **visualization**

33 Several computer programs were developed to meet specific research purpose in this study. All  
34 the programs are based on Igor Pro (www.wavemetrics.com) that provides a friendly GUI. Brief  
35 descriptions are given below.

36

### 37 **MRS program (Igor Pro based)**

38 The program (Figure S21) is written in Igor Pro (WaveMetrics, Inc. Lake Oswego, OR, USA)  
39 to feasible MRS calculation via a user-friendly GUI. The MRS application is not limited in  
40 SOC estimation, but can also be extended to other applications (e.g.  $E_{abs}$  estimation) as long as  
41 a reliable tracer is available.

42 MRS calculation can be done by different temporal cycles (batch calculation): by year, by  
43 year&season, by season, by year&month, by month, by year&month&hour. Data filter is also  
44 available to calculate MRS on a specific subset of data.

45 The program is available from <https://sites.google.com/site/wuchengust>.

46

### 47 **Mie program and source code written in Igor Pro**

48 A computer program (Figure S22) written in Igor Pro (WaveMetrics, Inc. Lake Oswego, OR,  
49 USA) for Mie scattering calculation. Both BHMIE and BHCOAT (coated particles)  
50 algorithms(Bohren and Huffman, 1983) are included. The program is also capable of batch  
51 calculation for both algorithms. Available from <https://sites.google.com/site/wuchengust>.

52

### 53 **Aethalometer data processing program (Igor Pro based)**

54 This handy tool (Figure S23) can perform different corrections (e.g. Weingartner, Virkkula) on  
55 Aethalometer data. Raw Aethalometer data suffers from several artifacts including filter matrix  
56 effect (multiple scattering), loading effect (shadowing) and scattering effect. Careful  
57 corrections are needed for reporting light absorption coefficient from attenuation measurement.  
58 This Igor based program can directly import Aethalometer raw data and perform corrections  
59 (algorithm can be selected by user). Results can be exported to .csv files. Extra information  
60 including statistics of sensor voltage from each channel, sampling flow rate, etc are plotted for  
61 a quick QA/QC check. Available from <https://sites.google.com/site/wuchengust>.

62

### 63 **Histbox program (Igor Pro based)**

64 A handy tool (Figure S24) to generate histogram and box plots with many powerful features.  
65 Data can be sorted by different time scale and batch plotting is available. Available from  
66 <https://sites.google.com/site/wuchengust>.

67

### 68 **Scatter plot program**

69 Scatter plot (Figure S25) is a handy tool to maximize the efficiency of data visualization in  
70 atmospheric science. The program includes Deming, WODR and York algorithm for linear  
71 regression, which consider uncertainties in both X and Y, that is more realistic for atmospheric  
72 applications. It is Igor based, and packed with lots of useful features for data analysis and graph  
73 plotting, including batch plotting, data masking via GUI, color coding in Z axis, data filtering  
74 and grouping. Available from <https://sites.google.com/site/wuchengust>.

75

76



## 77 Reference

- 78 Andreae, M. O., Schmid, O., Yang, H., Chand, D., Yu, J. Z., Zeng, L. M., and Zhang, Y. H.:  
79 Optical properties and chemical composition of the atmospheric aerosol in urban Guangzhou,  
80 China, *Atmos. Environ.*, 42, 6335-6350, doi: 10.1016/j.atmosenv.2008.01.030, 2008.
- 81 Bohren, C. F. and Huffman, D. R.: *Absorption and scattering of light by small particles*, Wiley,  
82 New York, xiv, 530 p. pp., 1983.
- 83 Chan, T. W., Brook, J. R., Smallwood, G. J., and Lu, G.: Time-resolved measurements of black  
84 carbon light absorption enhancement in urban and near-urban locations of southern Ontario,  
85 Canada, *Atmos. Chem. Phys.*, 11, 10407-10432, 2011.
- 86 Chow, J. C., Watson, J. G., Doraiswamy, P., Chen, L. W. A., Sodeman, D. A., Lowenthal, D.  
87 H., Park, K., Arnott, W. P., and Motallebi, N.: Aerosol light absorption, black carbon, and  
88 elemental carbon at the Fresno Supersite, California, *Atmos Res*, 93, 874-887, doi: DOI  
89 10.1016/j.atmosres.2009.04.010, 2009.
- 90 Chuang, P. Y., Duvall, R. M., Bae, M. S., Jefferson, A., Schauer, J. J., Yang, H., Yu, J. Z., and  
91 Kim, J.: Observations of elemental carbon and absorption during ACE-Asia and implications  
92 for aerosol radiative properties and climate forcing, *J. Geophys. Res.*, 108, 8634, doi: Doi  
93 10.1029/2002jd003254, 2003.
- 94 Doran, J. C., Barnard, J. C., Arnott, W. P., Cary, R., Coulter, R., Fast, J. D., Kassianov, E. I.,  
95 Kleinman, L., Laulainen, N. S., Martin, T., Paredes-Miranda, G., Pekour, M. S., Shaw, W. J.,  
96 Smith, D. F., Springston, S. R., and Yu, X. Y.: The T1-T2 study: evolution of aerosol properties  
97 downwind of Mexico City, *Atmos. Chem. Phys.*, 7, 1585-1598, doi: 10.5194/acp-7-1585-2007,  
98 2007.
- 99 Harris, D. C.: *Quantitative chemical analysis*, 8th ed., W.H. Freeman and Co., New York, 2010.
- 100 Knox, A., Evans, G. J., Brook, J. R., Yao, X., Jeong, C. H., Godri, K. J., Sabaliauskas, K., and  
101 Slowik, J. G.: Mass Absorption Cross-Section of Ambient Black Carbon Aerosol in Relation  
102 to Chemical Age, *Aerosol. Sci. Technol.*, 43, 522-532, doi: Doi 10.1080/02786820902777207,  
103 2009.
- 104 Lack, D. A. and Cappa, C. D.: Impact of brown and clear carbon on light absorption  
105 enhancement, single scatter albedo and absorption wavelength dependence of black carbon,  
106 *Atmos. Chem. Phys.*, 10, 4207-4220, doi: DOI 10.5194/acp-10-4207-2010, 2010.
- 107 Lan, Z.-J., Huang, X.-F., Yu, K.-Y., Sun, T.-L., Zeng, L.-W., and Hu, M.: Light absorption of  
108 black carbon aerosol and its enhancement by mixing state in an urban atmosphere in South  
109 China, *Atmos. Environ.*, 69, 118-123, doi: <http://dx.doi.org/10.1016/j.atmosenv.2012.12.009>,  
110 2013.
- 111 Liu, D., Flynn, M., Gysel, M., Targino, A., Crawford, I., Bower, K., Choularton, T., Jurányi,  
112 Z., Steinbacher, M., Hüglin, C., Curtius, J., Kampus, M., Petzold, A., Weingartner, E.,  
113 Baltensperger, U., and Coe, H.: Single particle characterization of black carbon aerosols at a  
114 tropospheric alpine site in Switzerland, *Atmos. Chem. Phys.*, 10, 7389-7407, doi: 10.5194/acp-  
115 10-7389-2010, 2010.
- 116 Mayol-Bracero, O. L., Gabriel, R., Andreae, M. O., Kirchstetter, T. W., Novakov, T., Ogren,  
117 J., Sheridan, P., and Streets, D. G.: Carbonaceous aerosols over the Indian Ocean during the  
118 Indian Ocean Experiment (INDOEX): Chemical characterization, optical properties, and  
119 probable sources, *J. Geophys. Res.*, 107, 8030, doi: Doi 10.1029/2000jd000039, 2002.
- 120 Moosmuller, H., Chakrabarty, R. K., Ehlers, K. M., and Arnott, W. P.: Absorption Angstrom  
121 coefficient, brown carbon, and aerosols: basic concepts, bulk matter, and spherical particles,  
122 *Atmos. Chem. Phys.*, 11, 1217-1225, doi: DOI 10.5194/acp-11-1217-2011, 2011.
- 123 Naoe, H., Hasegawa, S., Heintzenberg, J., Okada, K., Uchiyama, A., Zaizen, Y., Kobayashi, E.,  
124 and Yamazaki, A.: State of mixture of atmospheric submicrometer black carbon particles and

125 its effect on particulate light absorption, *Atmos. Environ.*, 43, 1296-1301, doi:  
126 <https://doi.org/10.1016/j.atmosenv.2008.11.031>, 2009.  
127 Pandolfi, M., Cusack, M., Alastuey, A., and Querol, X.: Variability of aerosol optical properties  
128 in the Western Mediterranean Basin, *Atmos. Chem. Phys.*, 11, 8189-8203, doi: DOI  
129 10.5194/acp-11-8189-2011, 2011.  
130 Thompson, J. E., Hayes, P. L., Jimenez, K. A. J. L., Zhang, X., Liu, J., Weber, R. J., and Buseck,  
131 P. R.: Aerosol Optical Properties at Pasadena, CA During CalNex 2010, *Atmos Environ*, doi:  
132 10.1016/j.atmosenv.2012.03.011, 2012.  
133 Wang, Q., Huang, R., Zhao, Z., Cao, J., Ni, H., Tie, X., Zhu, C., Shen, Z., Wang, M., and Dai,  
134 W.: Effects of photochemical oxidation on the mixing state and light absorption of black carbon  
135 in the urban atmosphere of China, *Environmental Research Letters*, 12, 044012, 2017.  
136 Wang, Q. Y., Huang, R. J., Cao, J. J., Han, Y. M., Wang, G. H., Li, G. H., Wang, Y. C., Dai,  
137 W. T., Zhang, R. J., and Zhou, Y. Q.: Mixing State of Black Carbon Aerosol in a Heavily  
138 Polluted Urban Area of China: Implications for Light Absorption Enhancement, *Aerosol. Sci.*  
139 *Technol.*, 48, 689-697, doi: 10.1080/02786826.2014.917758, 2014.  
140 Xu, J., Bergin, M. H., Yu, X., Liu, G., Zhao, J., Carrico, C. M., and Baumann, K.: Measurement  
141 of aerosol chemical, physical and radiative properties in the Yangtze delta region of China,  
142 *Atmos. Environ.*, 36, 161-173, 2002.  
143 Yang, M., Howell, S. G., Zhuang, J., and Huebert, B. J.: Attribution of aerosol light absorption  
144 to black carbon, brown carbon, and dust in China - interpretations of atmospheric measurements  
145 during EAST-AIRE, *Atmos. Chem. Phys.*, 9, 2035-2050, 2009.

146

147

**Table S1.** Comparison of Mass absorption efficiency (MAE) at various locations. For literature MAE values at different wavelengths rather than 550 nm, an estimated MAE<sub>550</sub> is given in the brackets following equations given by Moosmuller et al. (2011) assuming AAE of 1.

Location	Type	Sampling Duration	Inlet	$\lambda$ (nm)	$\sigma_{\text{abs}}$ Instrument	EC determination protocol	$\sigma_{\text{abs}} \pm 1$ S.D. (Mm <sup>-1</sup> )	EC mass ( $\mu\text{g m}^{-3}$ )	estimated MAE <sub>p</sub> * (m <sup>2</sup> g <sup>-1</sup> )	observed MAE (m <sup>2</sup> g <sup>-1</sup> )		Reference
										arithmetic mean $\pm 1$ S.D.	Gaussian fit	
Guangzhou, China	Suburban	2012.2-2013.1	PM <sub>2.5</sub>	550	AE	NIOSH_TOT	42.65±29.41	2.66±2.27	13*	18.75±6.16	16.16	This study
Shenzhen, China	Urban	2011.8-9	PM <sub>2.5</sub>	532	PAS	LII	25.4±19.0	4.0±3.1	/	6.5±0.5[6.29±0.48]	/	(Lan et al., 2013)
Xi'an, China	Urban	2012.12-2013.1	PM <sub>2.5</sub>	870	PAS	LII	/	8.8±7.3	7.17[11.34]	/	7.62[12.05]	(Wang et al., 2014)
Xi'an, China	Urban	2013.2	PM <sub>2.5</sub>	532	PAS	LII	/	/	/	14.6±5.6	12.7	(Wang et al., 2017)
Guangzhou, China	Urban	2004.10	PM <sub>2.5</sub>	532	PAS	NIOSH_TOT	91±60	7.1	7.7[7.44]	/	/	(Andreae et al., 2008)
Fresno, USA	Urban	2005.8-9	PM <sub>2.5</sub>	532	PAS	IMPROVE_A_TOR	5.06	1.01	/	6.1±2.5[5.9±2.42]	/	(Chow et al., 2009)
						NIOSH_TOT	/	0.58	/	9.3±2.4[8.99±2.32]	/	
T1, Mexico city, Mexico	Suburban	2006.3	PM <sub>2.5</sub>	870	PAS	NIOSH_TOT	/	/	/	9.2~9.7***[14.55~15.34]	/	(Doran et al., 2007)
Tokyo, Japan	Suburban	2005.8	PM <sub>2.5</sub>	565	PSAP	IMPROVE_A_TOR	30.43±20.41	2.9±2.13	11±1	/	/	(Naoue et al., 2009)
Pasadena, USA	Urban	2010.5-6	PM <sub>2.5</sub>	532	AM	NIOSH_TOT	3.8±3.4	0.6~0.7	5.7[5.51]	/	/	(Thompson et al., 2012)
Toronto, Canada	Urban	2006.12-2007.1	PM <sub>2.5</sub>	760	PAS	NIOSH_TOT	/	/	6.9~9.1** [9.53~12.57]	9.3~9.9[12.85~13.68]	/	(Knox et al., 2009)
Toronto, Canada	Suburban						3~6	0.10~0.14	/	30~43[42.6~61.06]	/	
Windsor, Canada	Urban	2007.8	PM <sub>2.5</sub>	781	PAS	LII	4.4±2.9	0.27±0.23	/	16±1[22.72±1.42]	/	(Chan et al., 2011)
Ottawa, Canada	Urban						26±17	1.7±0.9	/	15±3[21.3±4.26]	/	
Beijing, China	Rural	2005.3	/	550	AE	NIOSH_TOT	/	/	9.5	11.3	/	(Yang et al., 2009)
Montseny, Spin (Mediterranean)	Rural	2009.11-2010.10	PM <sub>10</sub>	637	MAAP	NIOSH_TOT	2.8±2.2	0.271±0.215	/	10.4[12.04]	/	(Pandolfi et al., 2011)
Jungfrauoch, Switzerland	Rural (high alpine)	2007.2-3	/	637	MAAP	LII	/	/	/	10.2±3.2[11.81±3.71]	/	(Liu et al., 2010)
Lin'an, China	Rural	1999.11	PM <sub>2.5</sub>	550	PSAP	NIOSH_TOT	23±14	3.4±1.7	/	8.6±7.0	/	(Xu et al., 2002)
Jeju Island, Korea	Coastal Rural, (East China Sea)	2001.4	PM <sub>10</sub>	550	PSAP	NIOSH_TOT	/	/	/	12.6±2.6	/	(Chuang et al., 2003)
Maldives	Oceanic rural	1999.2-3	PM <sub>3</sub>	550	PSAP	EGA	62±34	2.5±1.4	6.6	8.1	/	(Mayol-Bracero et al., 2002)

\*Determined by Minimum R Squared method; \*\* Median values;

AE: Aethalometer ; PAS photo acoustic spectrometer; MAAP: Multi Angle Absorption Photometer; PSAP: particle soot absorption photometer; AM: albedo meter; LII: Laser induced incandescence

**Table S2.** Statistics of monthly MAE<sub>550</sub>.

Month	95th	75th	50th	25th	5th	Mean	Max	Min	S.D.	N
Feb-2012	31.24	22.00	18.12	15.74	13.92	19.66	47.73	11.74	5.66	529
Mar-2012	26.51	19.63	17.45	15.91	13.71	18.46	45.56	10.98	4.30	651
Apr-2012	33.06	22.66	18.24	16.11	13.85	20.21	48.29	6.01	6.23	595
May-2012	33.24	23.25	19.59	17.16	14.82	21.07	46.66	6.33	5.62	528
Jun-2012	35.52	25.86	21.28	18.57	14.99	22.95	49.07	5.62	6.66	315
Jul-2012	33.93	23.77	18.81	15.58	12.71	20.51	49.22	9.23	6.79	587
Aug-2012	40.75	27.72	21.85	16.14	12.51	23.09	49.95	9.75	8.56	545
Sep-2012	30.75	20.86	17.52	15.24	12.97	18.99	46.44	10.39	5.63	674
Oct-2012	20.72	15.84	13.95	12.60	11.18	14.70	34.09	7.34	3.21	715
Nov-2012	26.45	18.10	15.43	13.70	11.89	16.75	39.32	8.34	4.72	495
Dec-2012	28.57	19.04	15.73	13.66	11.78	17.18	47.39	9.33	5.47	585
Jan-2013	21.53	16.24	14.47	13.03	11.80	15.29	43.19	7.16	3.77	708

**Table S3.** Statistics of monthly AAE<sub>470-660</sub>.

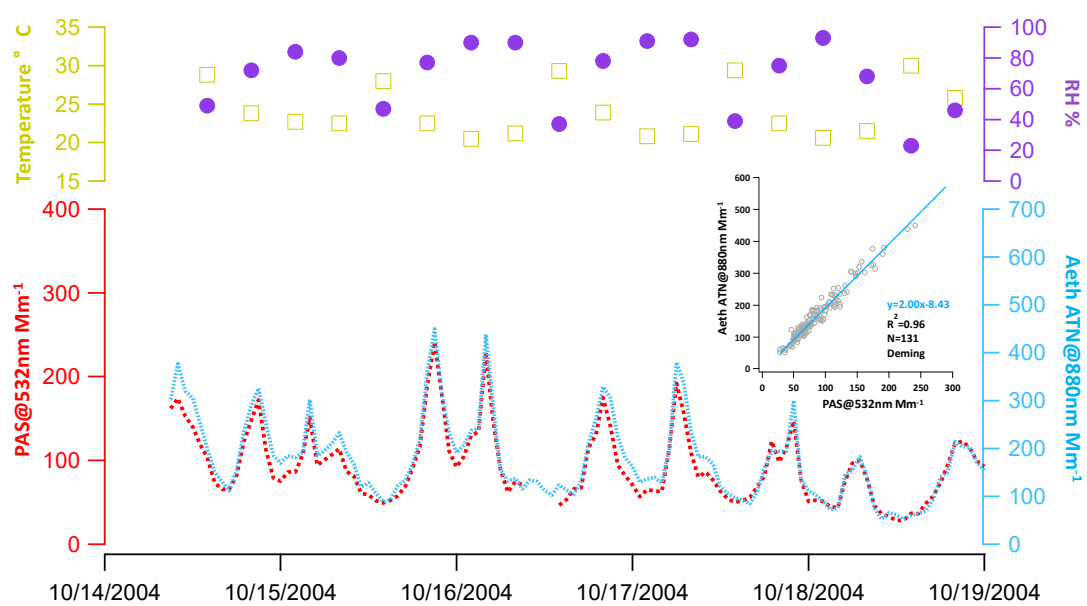
Month	95th	75th	50th	25th	5th	Mean	Max	Min	S.D.	N
Feb-2012	1.42	1.26	1.18	1.11	0.96	1.19	1.72	0.86	0.14	529
Mar-2012	1.33	1.18	1.10	1.02	0.93	1.10	1.50	0.65	0.12	651
Apr-2012	1.19	1.08	1.02	0.95	0.78	1.01	1.76	0.15	0.14	595
May-2012	1.13	1.05	1.00	0.94	0.84	0.99	1.24	0.39	0.10	528
Jun-2012	1.18	1.10	1.04	0.97	0.90	1.04	1.29	0.78	0.09	315
Jul-2012	1.22	1.11	1.04	0.98	0.83	1.04	1.43	0.20	0.13	587
Aug-2012	1.18	1.10	1.04	0.99	0.90	1.04	1.31	0.69	0.09	545
Sep-2012	1.23	1.13	1.06	1.00	0.91	1.07	1.40	0.64	0.11	674
Oct-2012	1.27	1.16	1.09	1.02	0.93	1.09	1.40	0.85	0.10	715
Nov-2012	1.24	1.16	1.11	1.05	0.97	1.11	1.52	0.79	0.08	495
Dec-2012	1.33	1.22	1.16	1.10	1.01	1.16	1.42	0.77	0.09	585
Jan-2013	1.38	1.27	1.18	1.12	1.01	1.19	1.66	0.93	0.11	708

**Table S4.** Statistics of monthly SSA<sub>525</sub>.

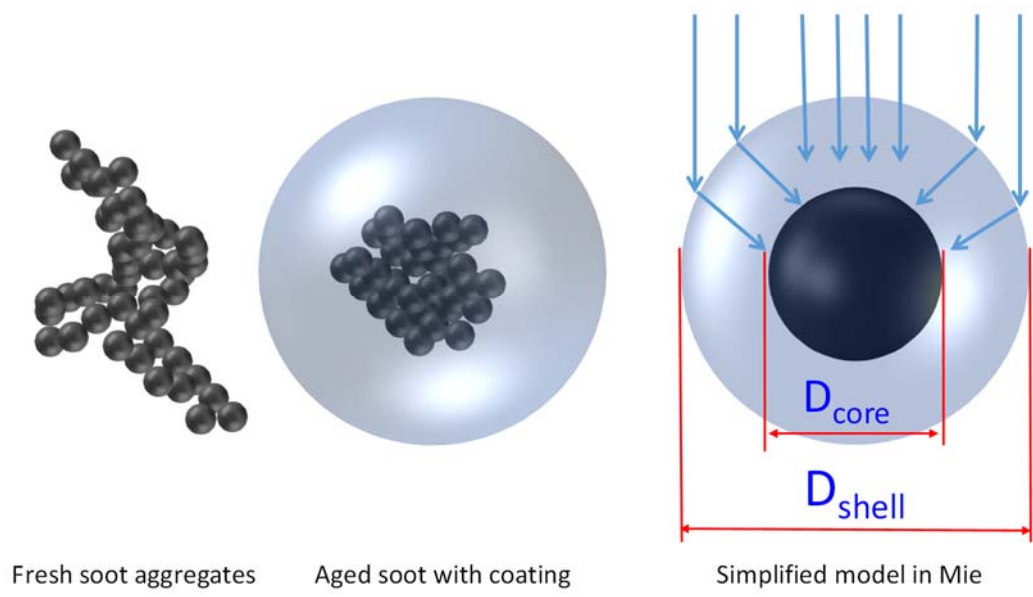
<b>Month</b>	<b>95th</b>	<b>75th</b>	<b>50th</b>	<b>25th</b>	<b>5th</b>	<b>Mean</b>	<b>Max</b>	<b>Min</b>	<b>S.D.</b>	<b>N</b>
Feb-2012	0.91	0.89	0.87	0.84	0.79	0.86	0.94	0.65	0.04	526
Mar-2012	0.91	0.89	0.86	0.83	0.77	0.85	0.95	0.42	0.05	648
Apr-2012	0.92	0.89	0.86	0.83	0.76	0.85	0.94	0.45	0.06	552
May-2012	0.92	0.90	0.87	0.83	0.74	0.85	0.94	0.45	0.06	527
Jun-2012	0.92	0.89	0.86	0.81	0.74	0.84	0.95	0.64	0.06	310
Jul-2012	0.91	0.87	0.83	0.79	0.71	0.83	0.95	0.57	0.06	580
Aug-2012	0.94	0.92	0.89	0.85	0.79	0.88	0.96	0.67	0.05	536
Sep-2012	0.93	0.91	0.88	0.84	0.75	0.87	0.96	0.55	0.06	672
Oct-2012	0.94	0.93	0.91	0.89	0.84	0.90	0.96	0.66	0.03	715
Nov-2012	0.91	0.90	0.87	0.83	0.75	0.86	0.94	0.18	0.06	495
Dec-2012	0.91	0.89	0.86	0.82	0.74	0.85	0.94	0.66	0.05	585
Jan-2013	0.91	0.89	0.87	0.85	0.80	0.86	0.93	0.64	0.04	708

**Table S5.** Statistics of monthly E<sub>abs550</sub>.

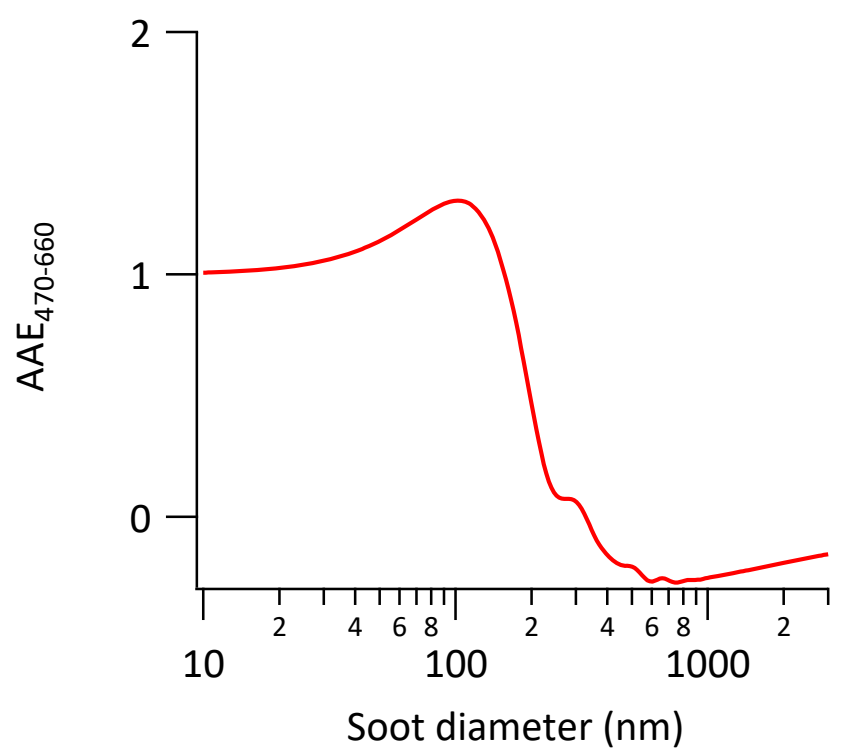
<b>Month</b>	<b>95th</b>	<b>75th</b>	<b>50th</b>	<b>25th</b>	<b>5th</b>	<b>Mean</b>	<b>Max</b>	<b>Min</b>	<b>S.D.</b>	<b>N</b>
Feb-2012	2.23	1.59	1.31	1.14	1.04	1.43	3.41	1.00	0.40	501
Mar-2012	1.76	1.30	1.15	1.07	1.02	1.24	2.83	1.00	0.26	466
Apr-2012	2.46	1.70	1.37	1.21	1.07	1.52	3.58	1.00	0.45	576
May-2012	2.48	1.73	1.45	1.28	1.12	1.57	3.46	1.02	0.41	520
Jun-2012	2.47	1.80	1.49	1.30	1.09	1.61	3.38	1.01	0.45	305
Jul-2012	2.83	1.97	1.57	1.30	1.10	1.71	4.03	1.01	0.55	568
Aug-2012	3.45	2.34	1.86	1.40	1.11	1.97	4.20	1.00	0.71	528
Sep-2012	2.36	1.62	1.36	1.20	1.05	1.48	3.54	1.00	0.42	636
Oct-2012	1.87	1.44	1.27	1.15	1.04	1.34	3.07	1.00	0.28	683
Nov-2012	2.14	1.51	1.27	1.13	1.04	1.38	3.17	1.00	0.38	461
Dec-2012	2.46	1.64	1.35	1.18	1.05	1.49	4.02	1.00	0.46	555
Jan-2013	1.85	1.39	1.24	1.12	1.03	1.31	3.66	1.00	0.32	672



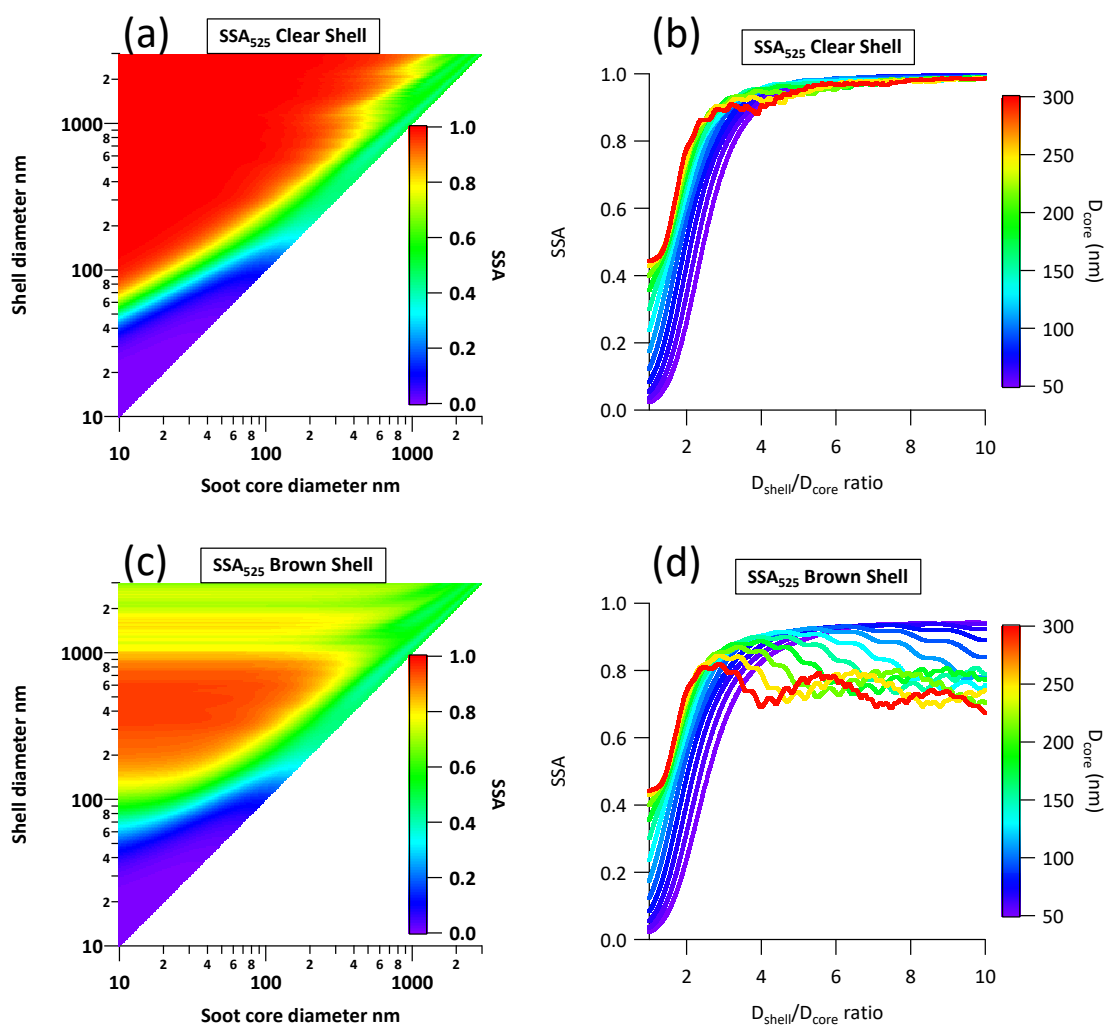
**Figure S1.** Comparison of collocated Aethalometer and PAS at Guangzhou (Oct 2004). Both PAS and Aethalometer (AE-16) were equipped with PM<sub>2.5</sub> inlets. RH of the sampled air was controlled to be <45% for PAS. Aethalometer sampling was conducted without RH control.



**Figure S2.** Schematic of the aging effect on light absorption. More light is absorbed by the soot particle core due to the lensing effect of the coating materials.

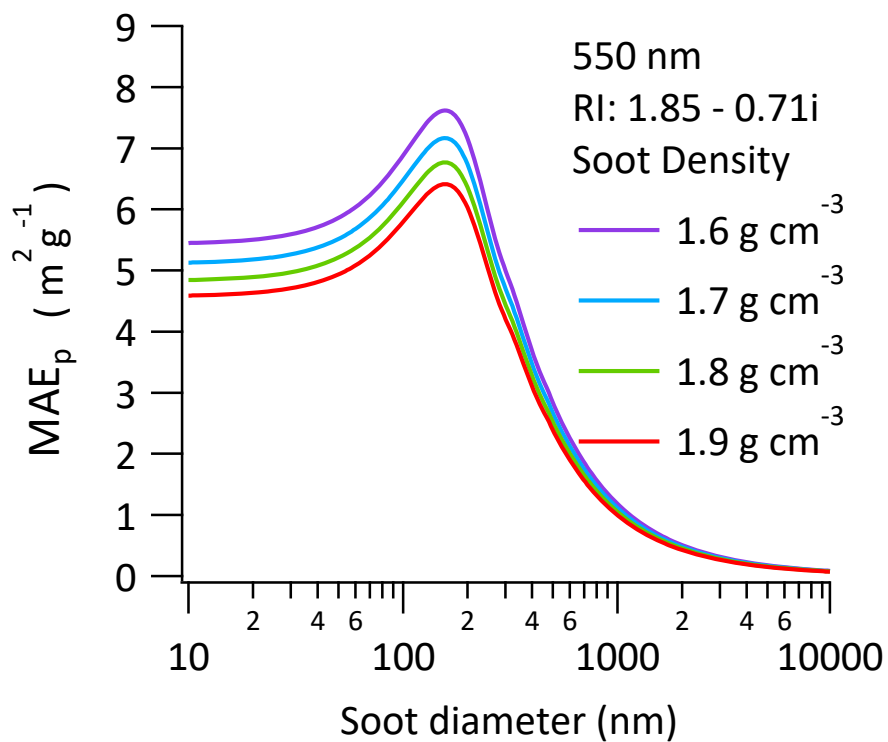


**Figure S3.** Mie simulated AAE<sub>470-660</sub> of a bare soot particle as a function of diameter with a Refractive index of 1.85 – 0.71i.

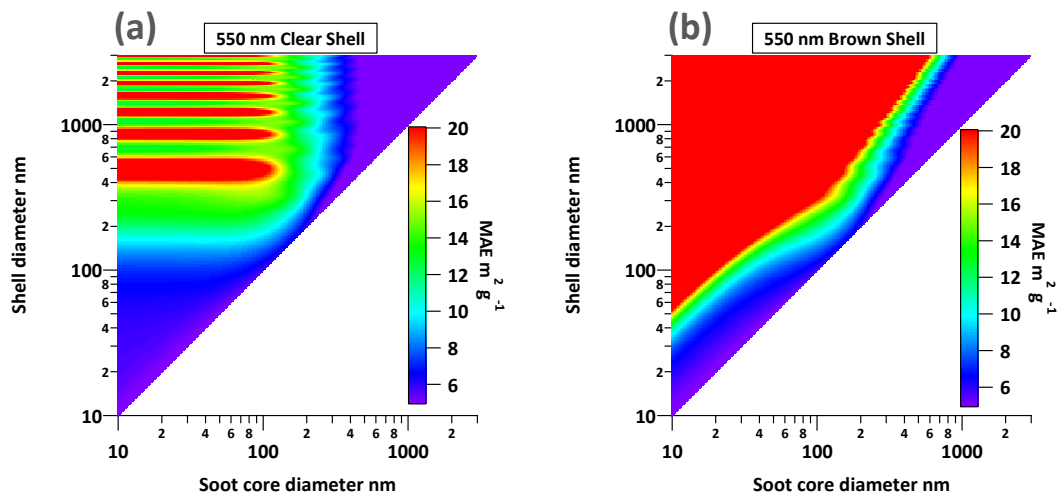


**Figure S4.** Mie simulated size dependency of soot particles SSA at wavelength 525 nm. (a) Combination of different clear shell (y axis) and core diameters (x axis). The color coding represents the SSA of a particle with specific core and clear shell size; (b) Cross-sections views of (a). The color coding represents different  $D_{\text{core}}$  in the range of 50 – 300 nm. (c)&(d) Similar to (a)&(b) but from the brown shell scenario.

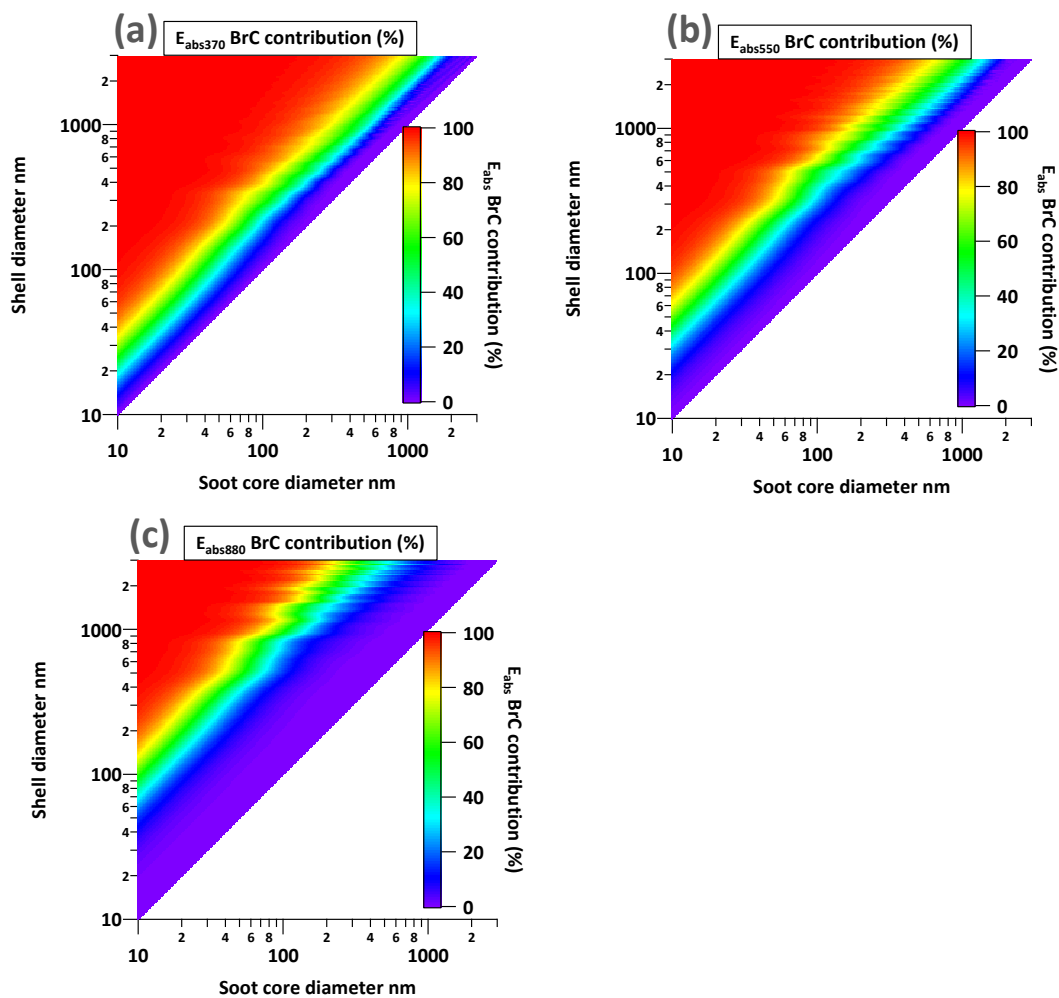




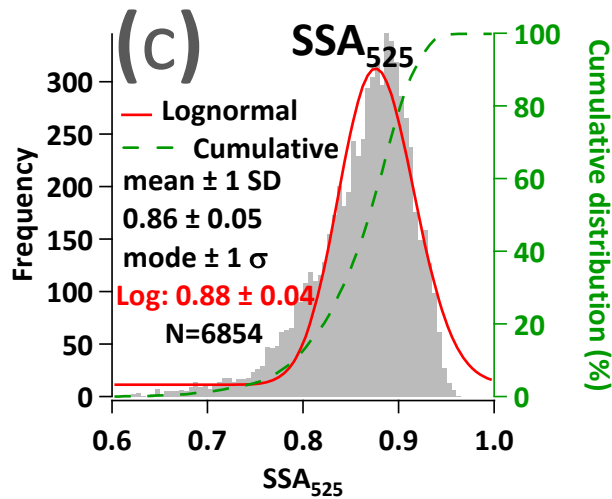
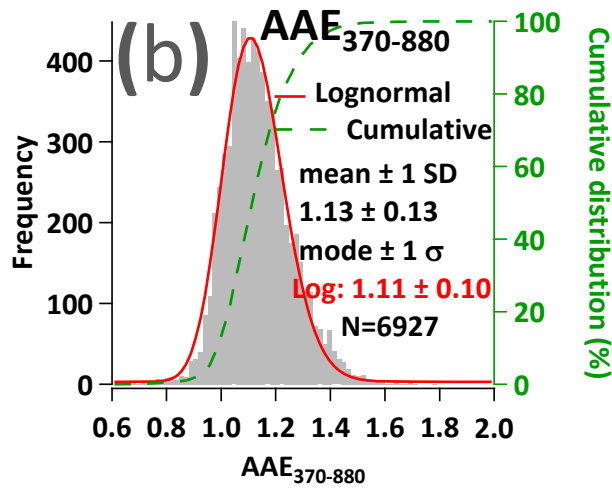
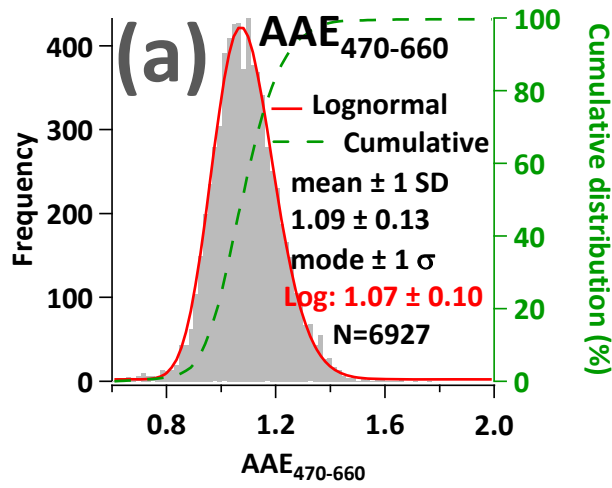
**Figure S5.** Mie simulated mass absorption efficiency (MAE<sub>p</sub>) of a bare soot particle as a function of diameter at a wavelength of 550nm. Refractive index is 1.85 – 0.71i and density varied from 1.6 to 1.9 g cm<sup>-3</sup>.



**Figure S6.** Mie simulated mass absorption efficiency (MAE) of a bare soot particle as a function of diameter at a wavelength of 550nm. Refractive index is  $1.85 - 0.71i$  and density is  $1.9 \text{ g cm}^{-3}$  for the soot core. Refractive index for clear coating is 1.55. Refractive index for brown coating is wavelength dependent adopted from Lack and Cappa (2010).



**Figure S7.** Mie simulated BrC absorption contribution to total  $E_{abs}$  (lensing effect + BrC absorption) in the brown shell scenario. (a) 370 nm (b) 550 nm (c) 880 nm.

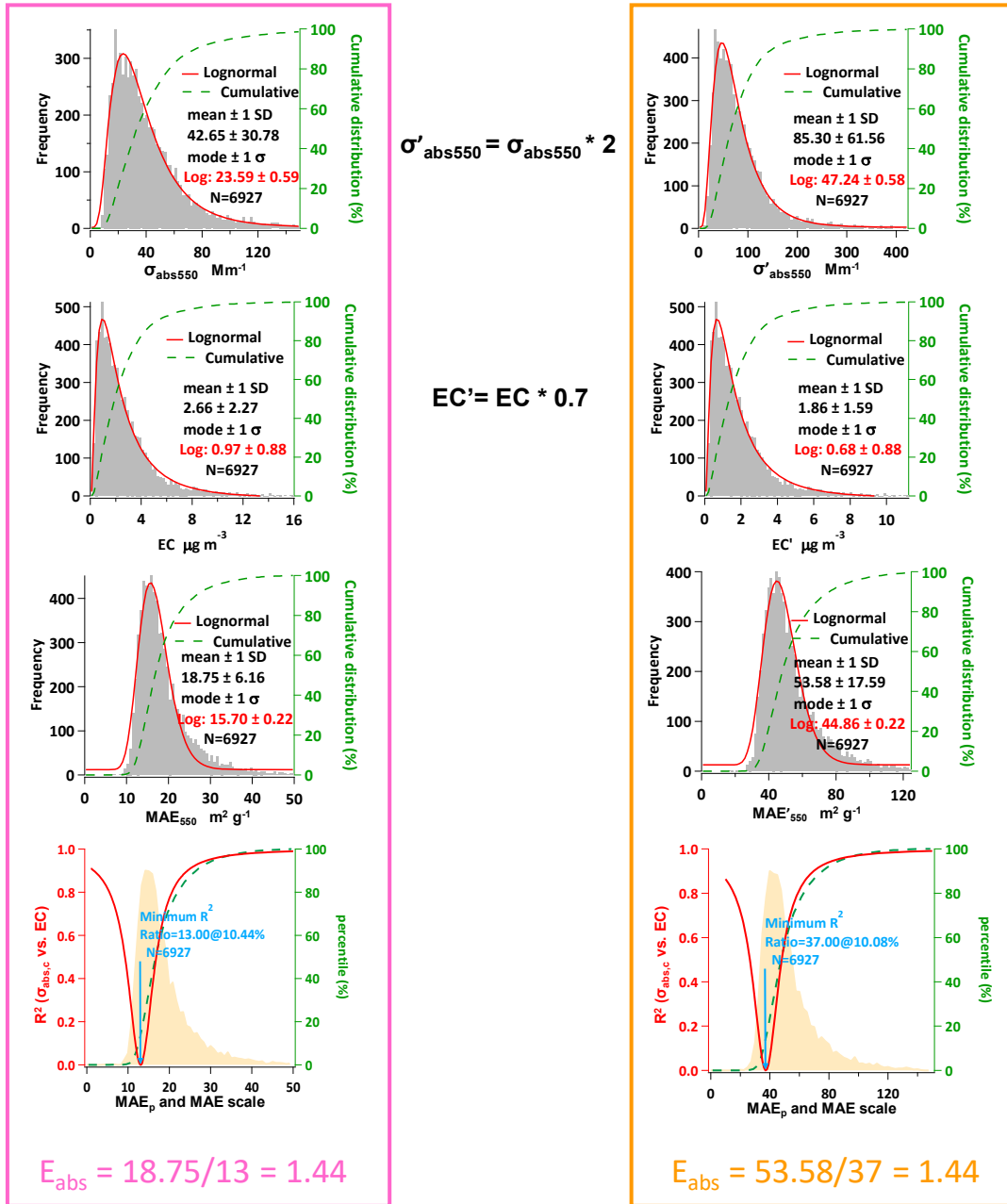


**Figure S8.** Measured annual statistics of AAE and SSA. (a) Annual frequency distribution of AAE<sub>470-660</sub>. (b) Annual frequency distribution of AAE<sub>370-880</sub>. (c) Annual frequency distribution of SSA<sub>525</sub>. The red line represents lognormal fitting curve.

# Test A

Original data

Systematically biased data

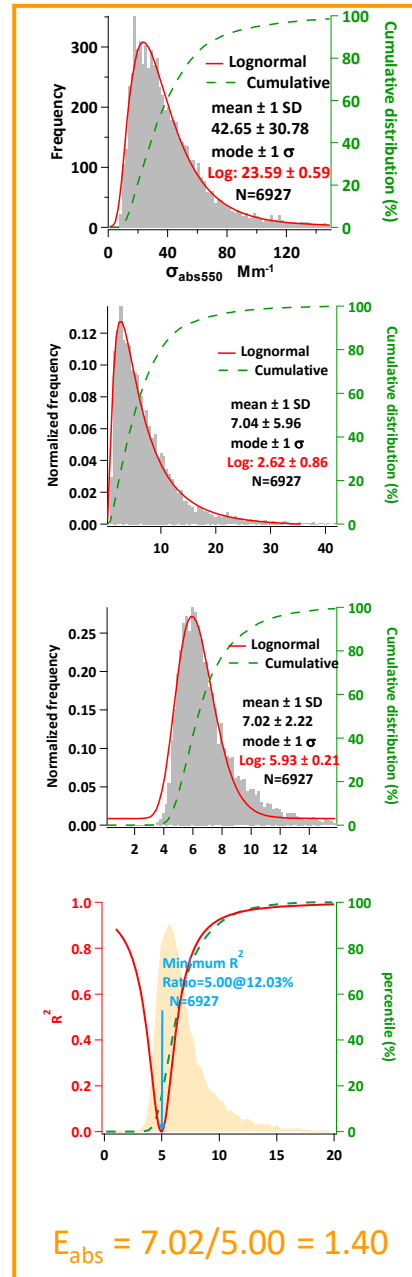
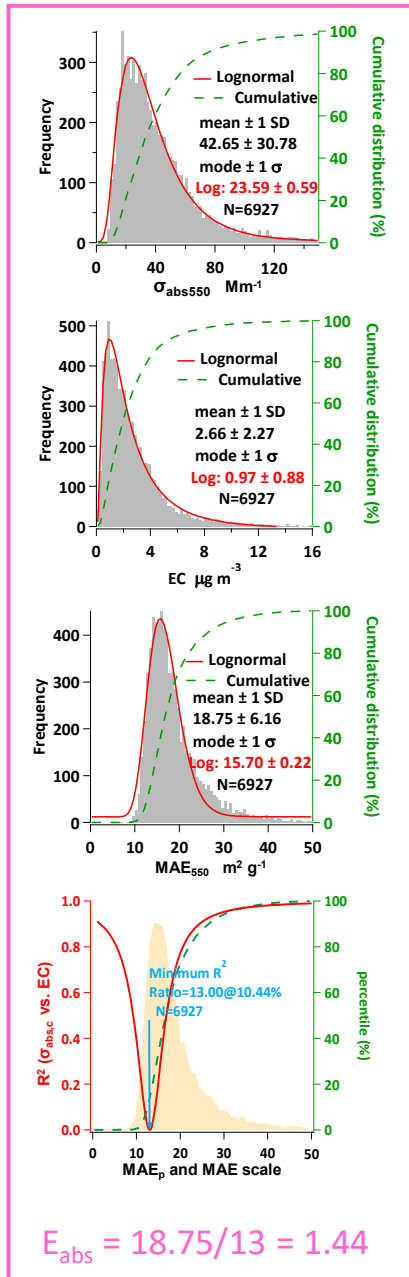


**Figure S9.** Comparison of  $E_{abs}$  from original data and systematically biased data (Test A). It should be noted that the  $E_{abs}$  shown here is ratio of averages, which is different from the annual average  $E_{abs}$  calculated from average of ratios.

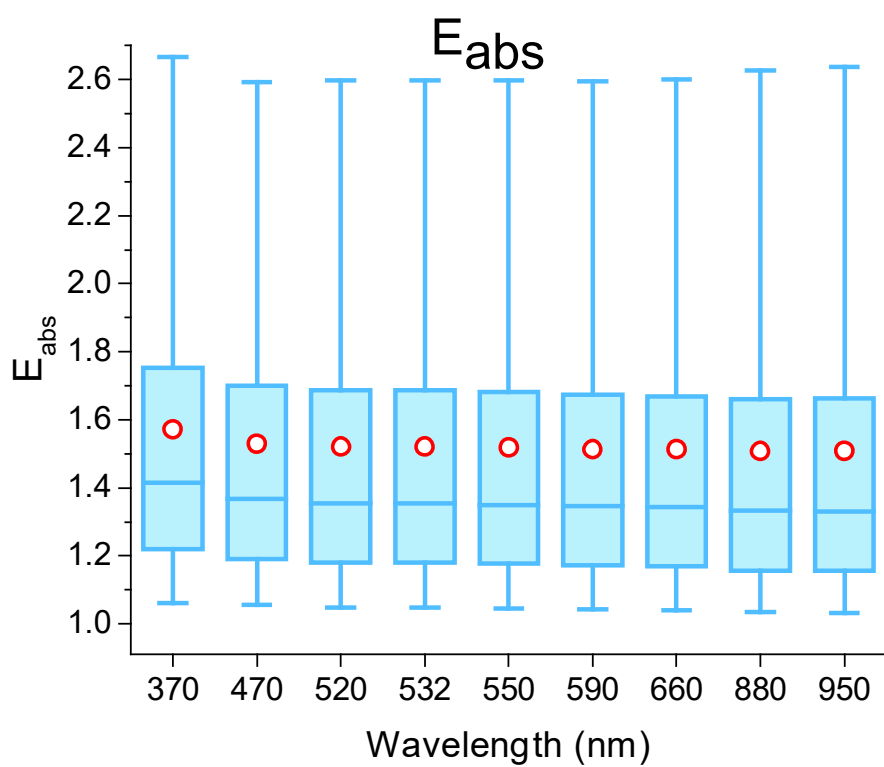
# Test B

Original data

Systematically biased data



**Figure S10.** Comparison of  $E_{\text{abs}}$  from data using NIOSH EC and data using IMPROVE EC (Test B). It should be noted that the  $E_{\text{abs}}$  shown here is ratio of averages, which is different from the annual average  $E_{\text{abs}}$  calculated from average of ratios.



Wavelength (nm)	370	470	520	532	550	590	660	880	950
$E_{\text{abs}}$ mean	1.55	1.51	1.50	1.50	1.50	1.49	1.49	1.48	1.49
$E_{\text{abs}}$ S.D.	0.48	0.47	0.47	0.47	0.48	0.47	0.48	0.49	0.49

**Figure S11.** Spectrum annual average  $E_{\text{abs}}$  from 370 to 950 nm.

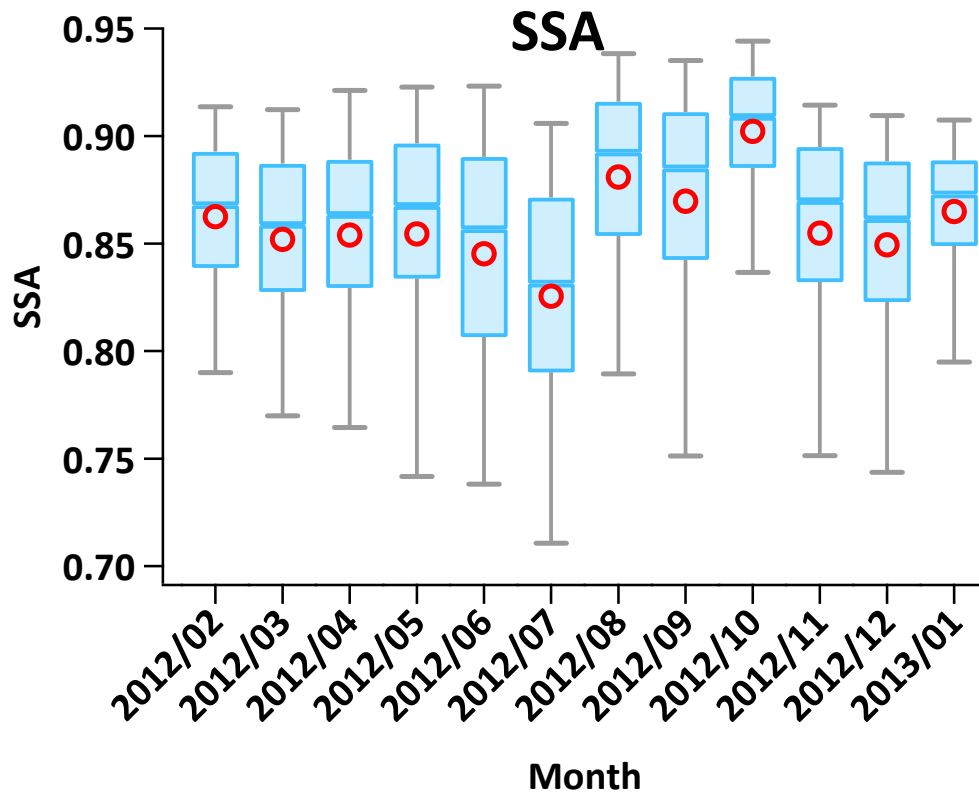
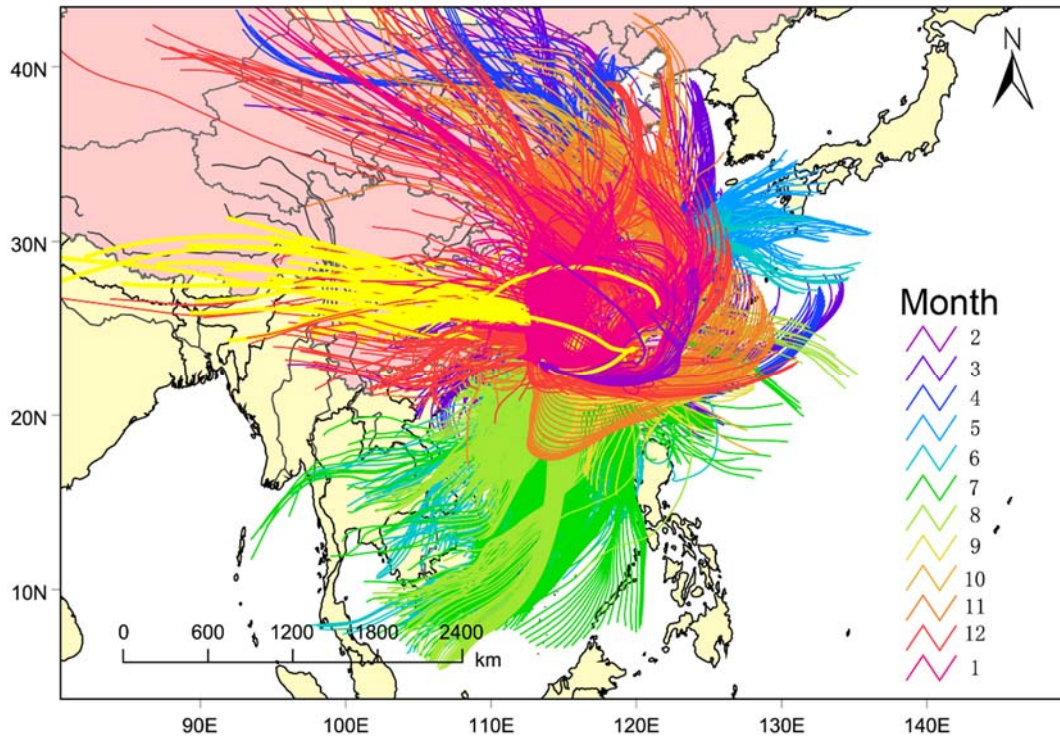
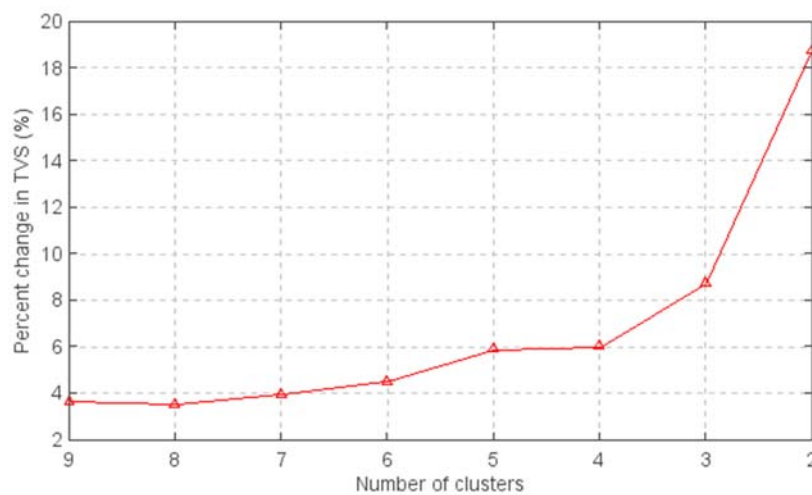


Figure S12. Measured monthly variations of SSA<sub>525</sub>.

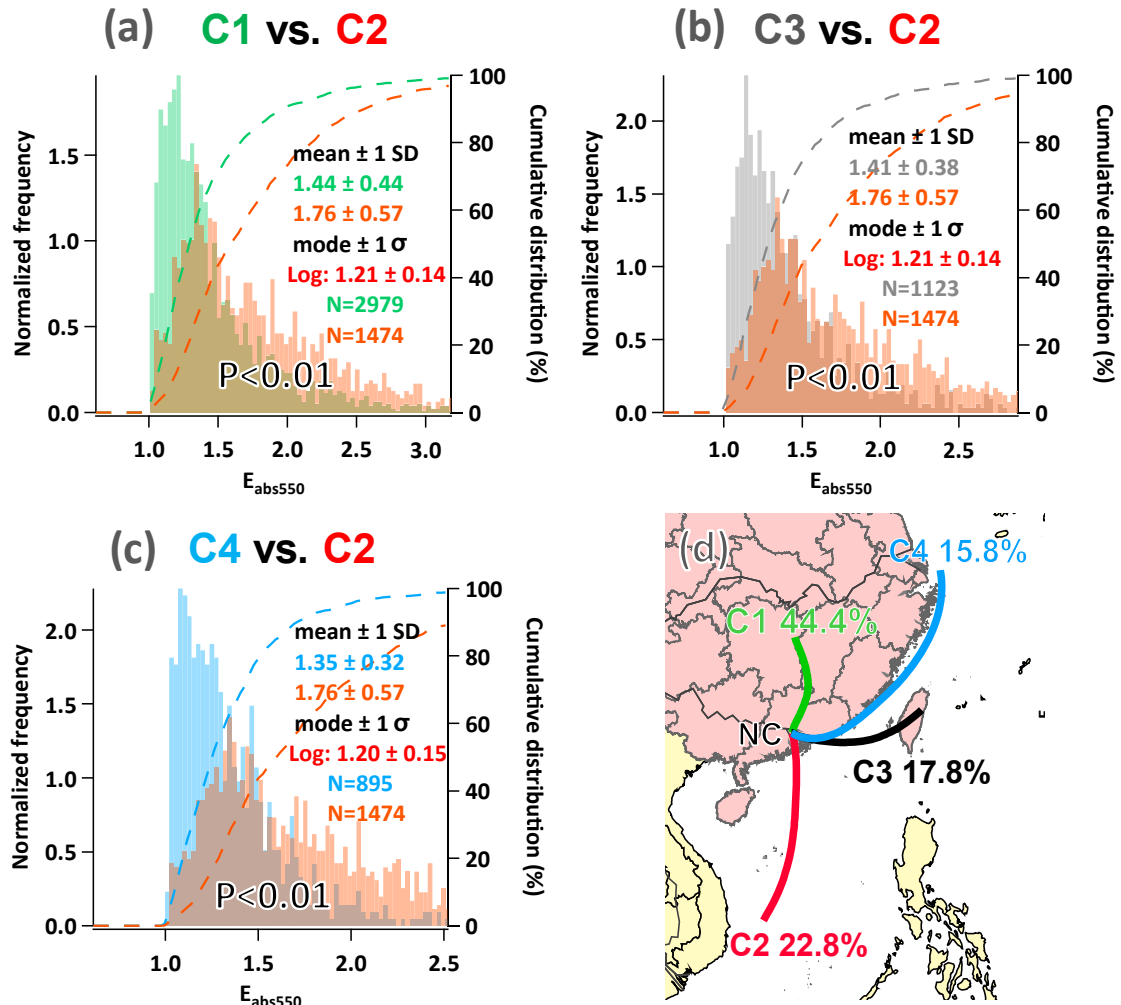




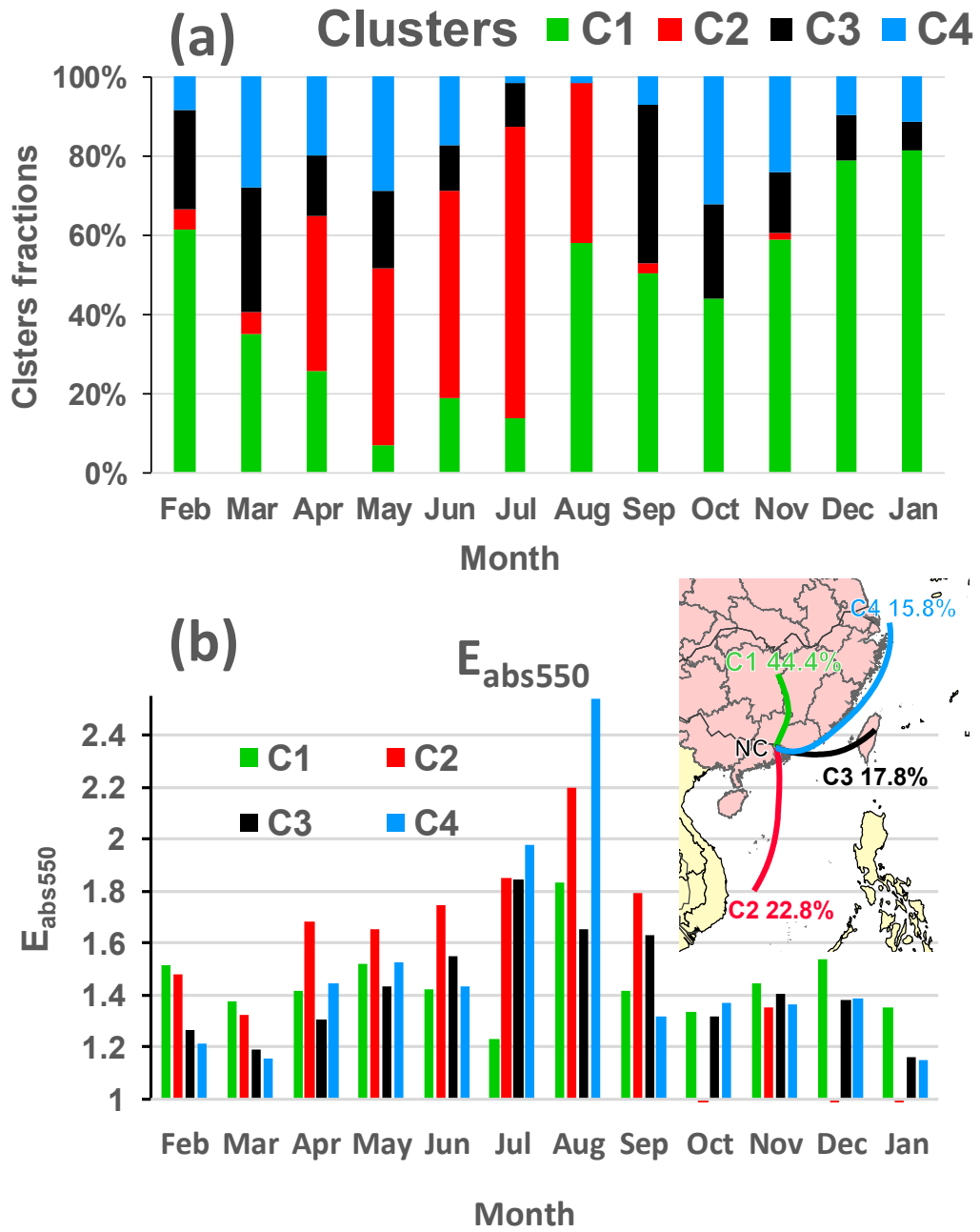
**Figure S13.** Hourly back trajectories for the past 72 hours calculated using NOAA's HYSPLIT model from Feb 2012 to Jan 2013. The color coding represents different months.



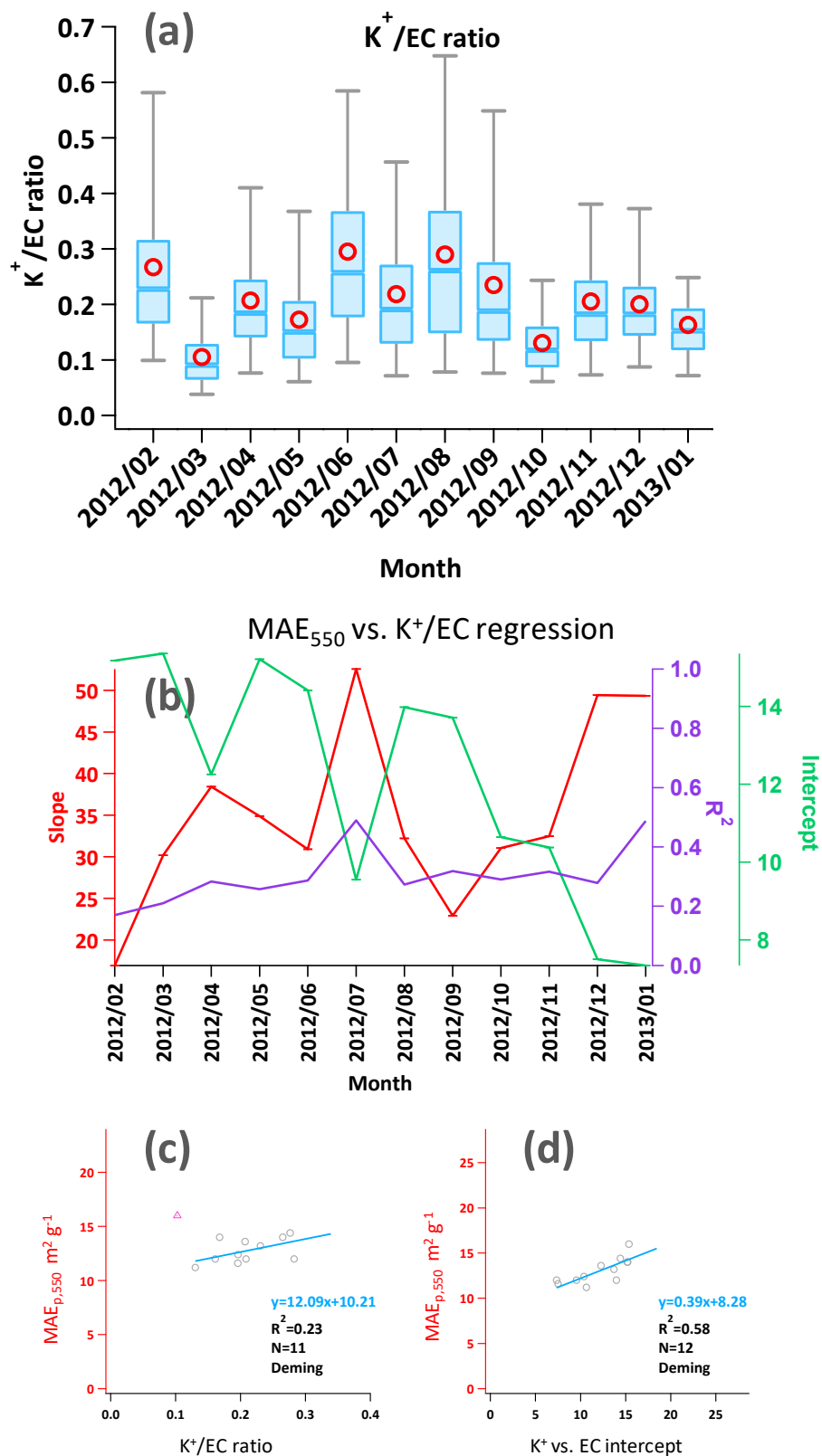
**Figure S14.** Total spatial variance (TSV) as a function of number of clusters in back trajectories clustering analysis.



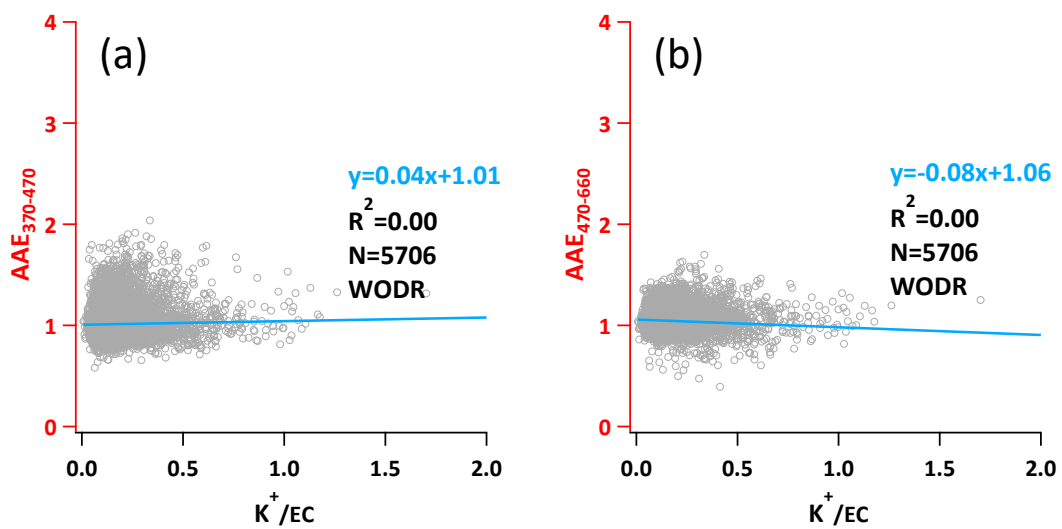
**Figure S15.** Frequency distributions of  $E_{abs550}$  by different air mass clusters. P is calculated by Wilcoxon-Mann-Whitney tests.



**Figure S16.** (a) Monthly contribution of each cluster. (b) Monthly  $E_{abs550}$  of each cluster.



**Figure S17.** (a) Monthly variations of K<sup>+</sup>/EC ratio from 2012 Feb to 2013 Jan at NC site. (b) Monthly regressions between MAE<sub>550</sub> and K<sup>+</sup>/EC with slope in red, intercept in green and R<sup>2</sup> in purple. (c) regressions between monthly MAE<sub>p,550</sub> and K<sup>+</sup>/EC. (d) regression between monthly MAE<sub>p,550</sub> and intercepts from (b).



**Figure S18.** Correlations of AAE with  $K^+/EC$  ratio (biomass burning indicator). (a) AAE from 370 – 470 nm. (b) AAE from 470 – 660 nm.

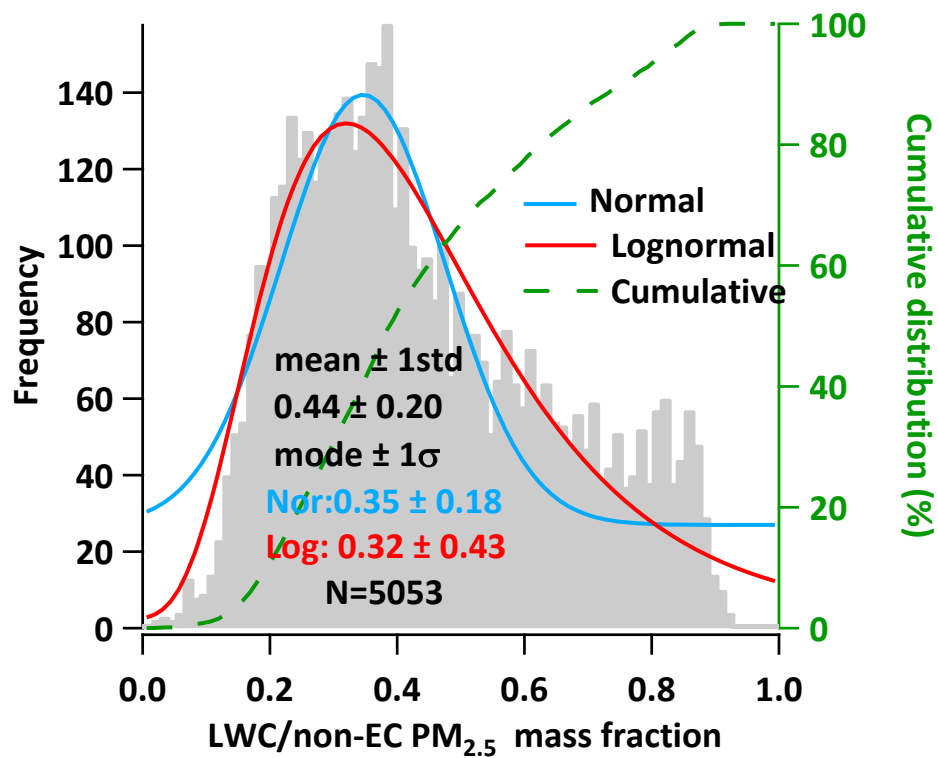
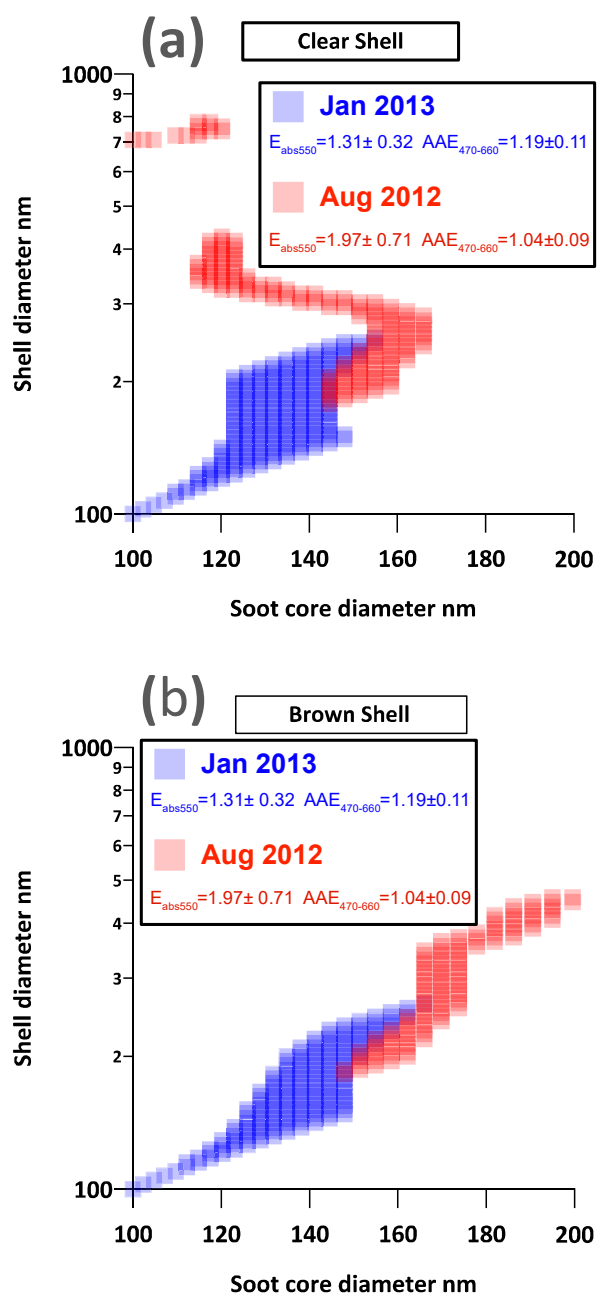


Figure S19. Annual frequency distribution of LWC/non-EC PM<sub>2.5</sub> mass fraction.



**Figure S20.** Size range of soot particles constrained by  $E_{\text{abs}}$ ,  $SSA_{525}$  and  $AAE_{470-660}$  from measurements. (a) Clear shell scenario; (b) Brown shell scenario

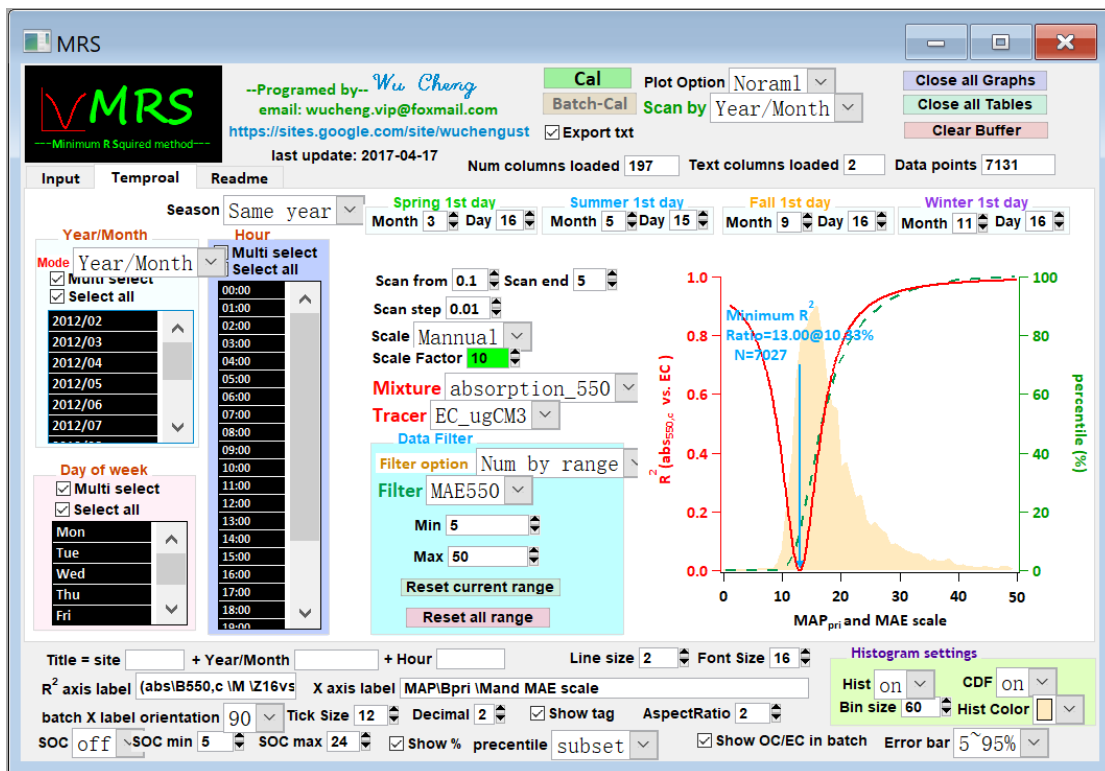
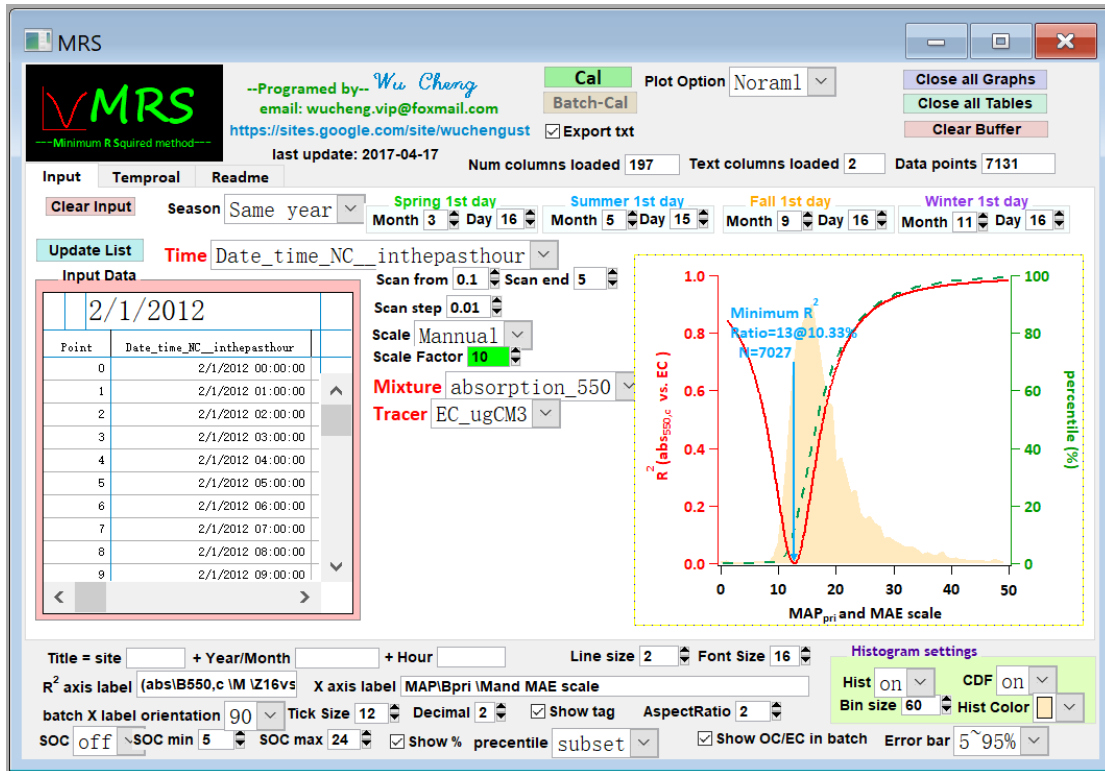
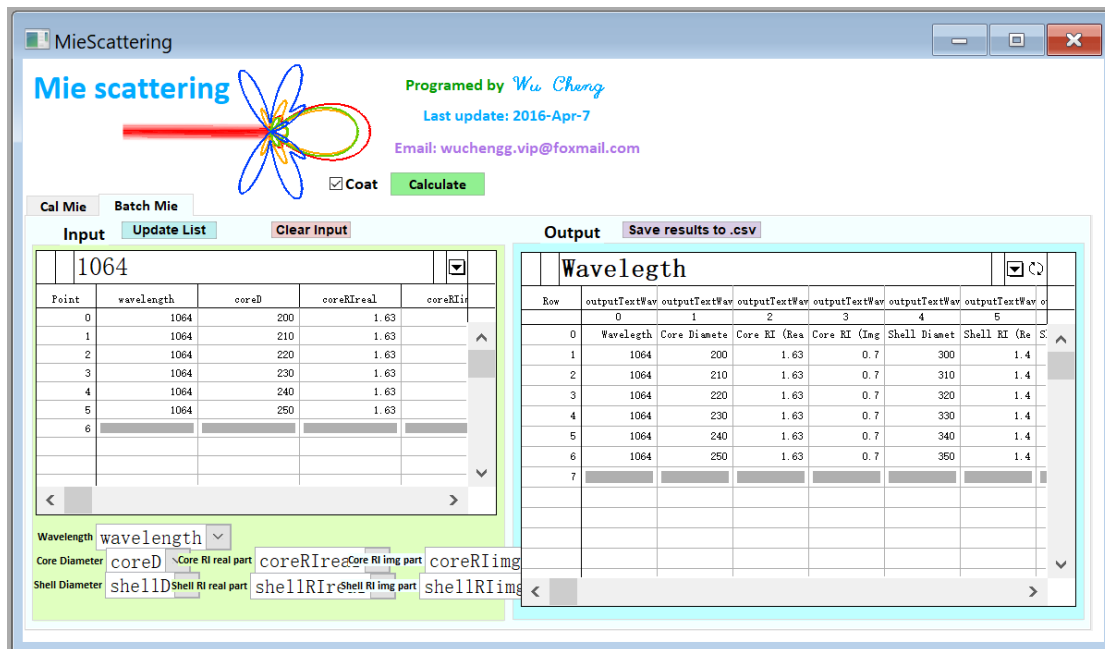
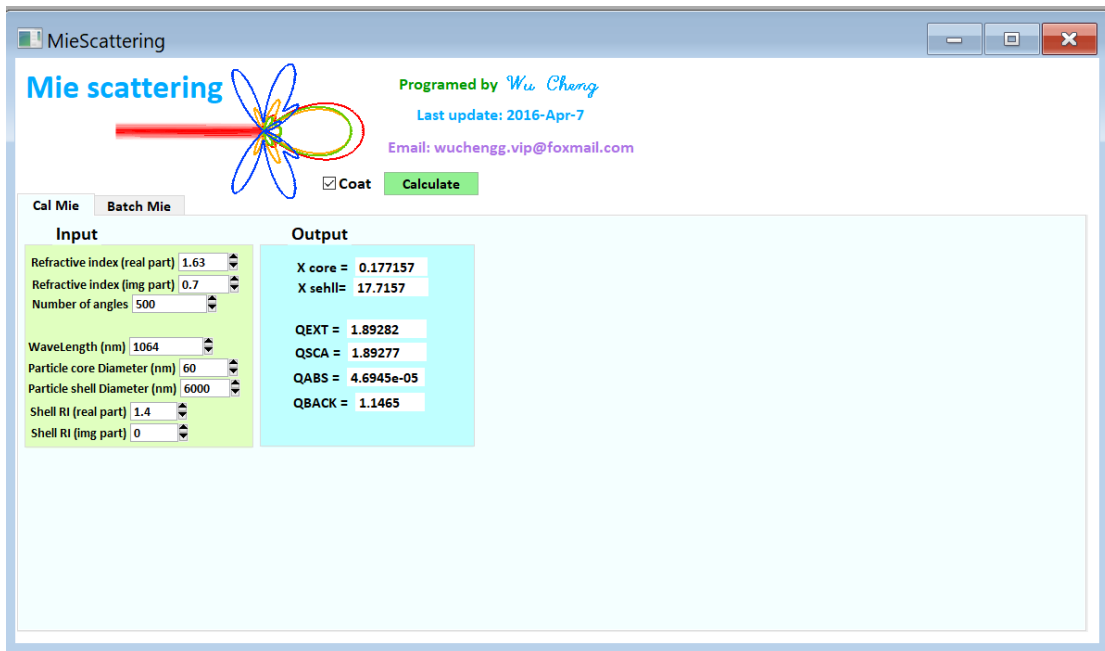


Figure S21. MRS program written in Igr Pro (WaveMetrics, Inc. Lake Oswego, OR, USA). Available from <https://sites.google.com/site/wuchengust>.





**Figure S22.** Mie program written in Igro Pro (WaveMetrics, Inc. Lake Oswego, OR, USA). Available from <https://sites.google.com/site/wuchengust>.

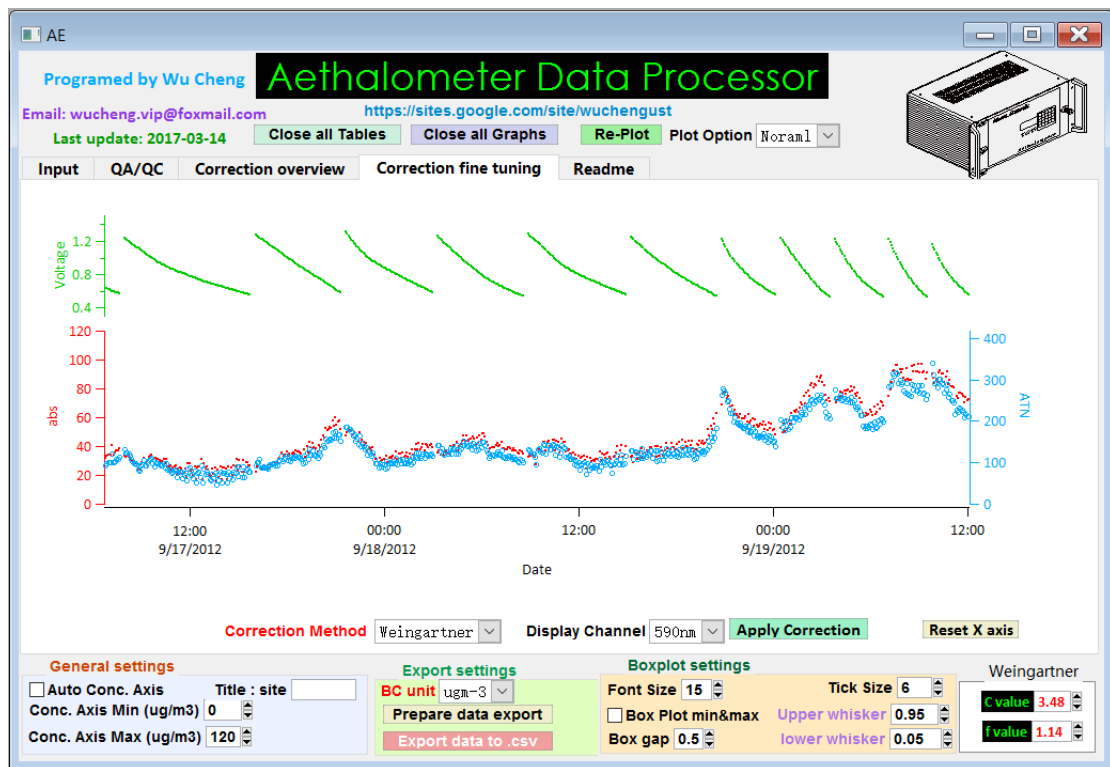
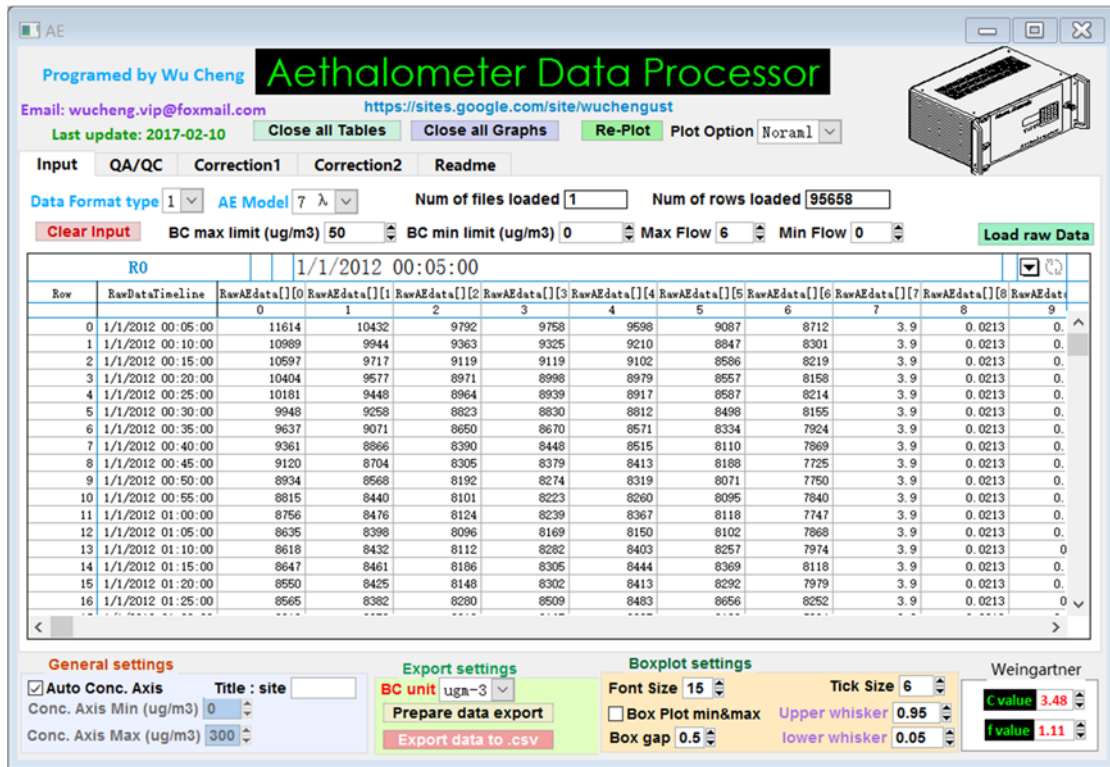


Figure S23. Aethalometer data processing program written in Igro Pro (WaveMetrics, Inc. Lake Oswego, OR, USA). Available from <https://sites.google.com/site/wuchengust>.

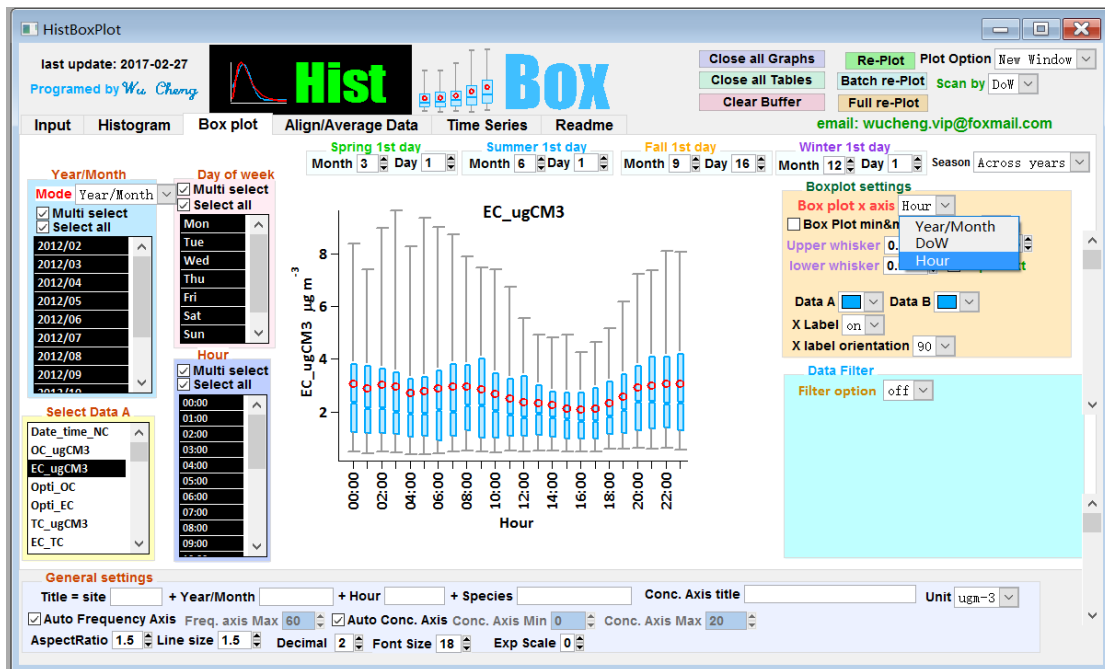
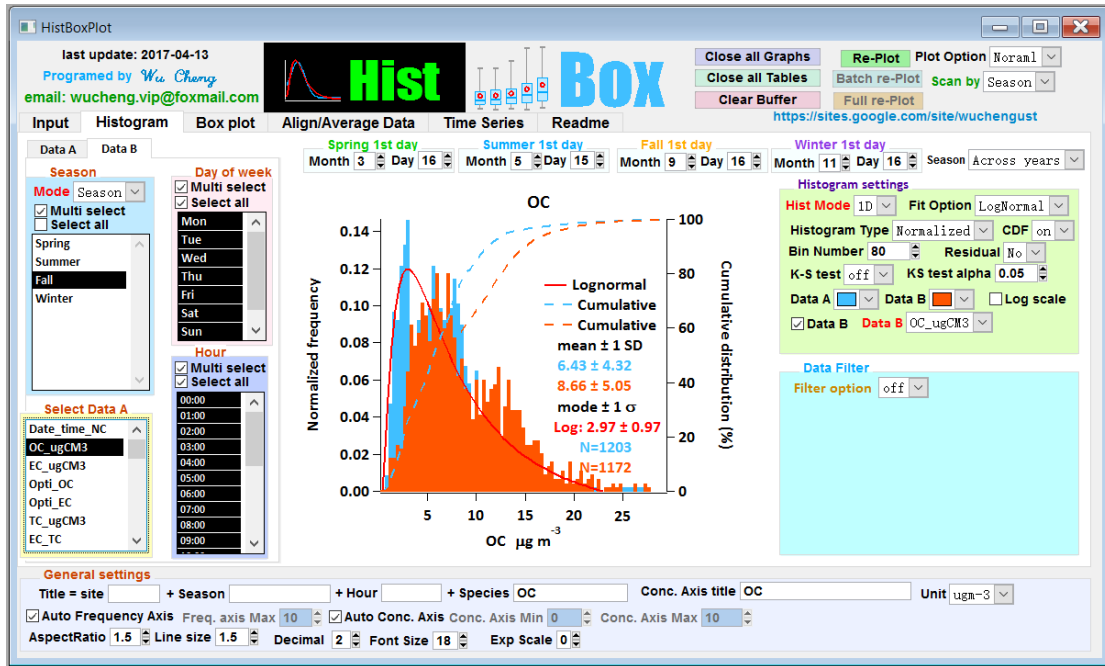


Figure S24. Histbox program written in Igro Pro (WaveMetrics, Inc. Lake Oswego, OR, USA). Available from <https://sites.google.com/site/wuchengust>.

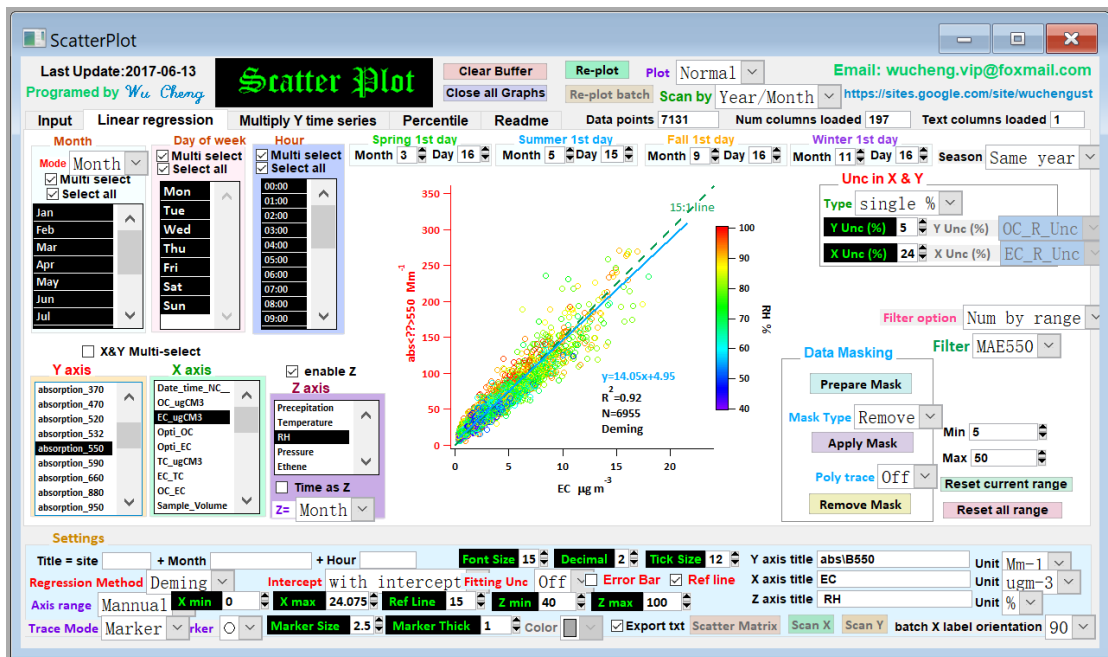
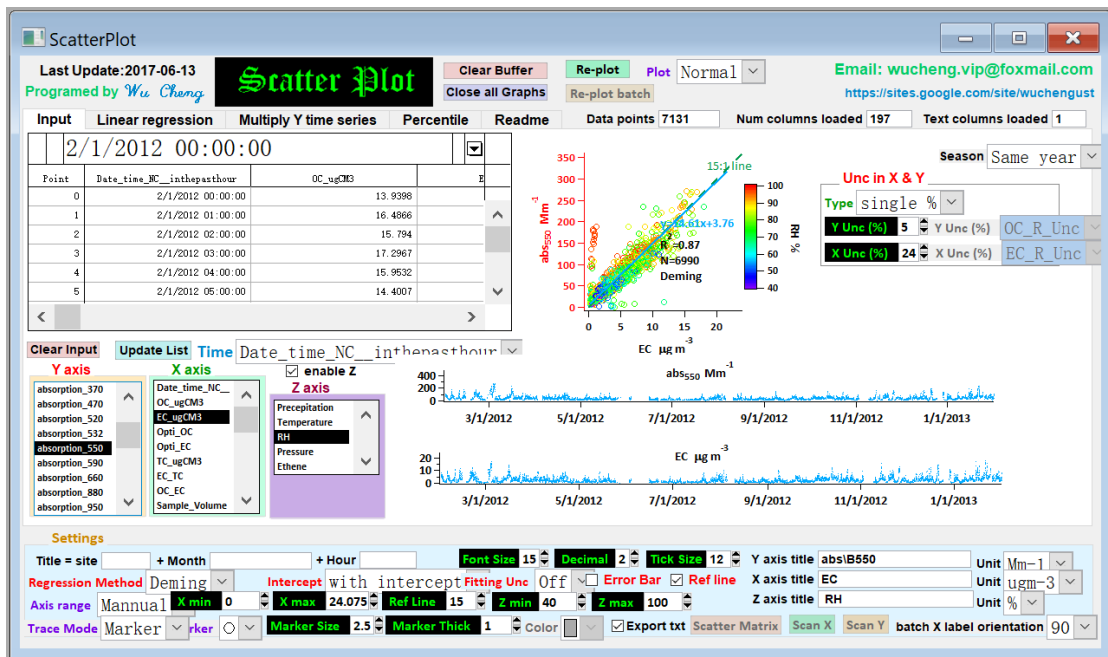


Figure S25. Scatter plot program written in Igro Pro (WaveMetrics, Inc. Lake Oswego, OR, USA). Available from <https://sites.google.com/site/wuchengust>.