

# Long-term profiling of mineral dust and pollution aerosol with multiwavelength polarization/Raman lidar at the Central Asian site of Dushanbe, Tajikistan: Case studies

Julian Hofer<sup>1</sup>, Dietrich Althausen<sup>1</sup>, Sabur F. Abdullaev<sup>2</sup>, Abduvosit N. Makhmudov<sup>2</sup>, Bakhron I. Nazarov<sup>2</sup>, Georg Schettler<sup>3</sup>, Ronny Engelmann<sup>1</sup>, Holger Baars<sup>1</sup>, K. Wadinga Fomba<sup>1</sup>, Konrad Müller<sup>1</sup>, Bernd Heinold<sup>1</sup>, Konrad Kandler<sup>4</sup>, and Albert Ansmann<sup>1</sup>

<sup>1</sup>Leibniz Institute for Tropospheric Research, Leipzig, Germany

<sup>2</sup>Physical Technical Institute of the Academy of Sciences of Tajikistan, Dushanbe, Tajikistan

<sup>3</sup>Helmholtz Center Potsdam, German Research Center for Geosciences, Potsdam, Germany

<sup>4</sup>Institut für Angewandte Geowissenschaften, Technische Universität Darmstadt, Darmstadt, Germany

*Correspondence to:* Julian Hofer, hofer@tropos.de

**Abstract.** For the first time, continuous vertically resolved aerosol measurements were performed by lidar in Tajikistan, Central Asia. Observations with the multiwavelength polarization/Raman lidar Polly<sup>XT</sup> were conducted during CADEX (Central Asian Dust EXperiment) in Dushanbe, Tajikistan, from March 2015 to August 2016. Co-located with the lidar a sun photometer was operated. The goal of CADEX is to provide an unprecedented data set on vertically resolved aerosol optical properties in Central Asia, an area highly affected by climate change but largely missing vertically resolved aerosol measurements. During the 18-months measurement campaign, mineral dust was detected frequently from ground to cirrus level height. In this study, an overview of the measurement period is given and four typical but different example measurement cases are discussed in detail. Three of them are dust cases and one is a contrasting pollution aerosol case. Vertical profiles of the measured optical properties and the calculated dust and non-dust mass concentrations are presented. Dust source regions were identified by means of backward trajectory analyses. A lofted layer of Middle Eastern dust with an aerosol optical thickness (AOT) of 0.4 and an extinction-related Ångström exponent of 0.41 was measured. In comparison, two near-ground dust cases have Central Asian sources. One is an extreme dust event with an AOT of 1.5 and Ångström exponent of 0.12 and the other one is a most extreme dust event with an AOT of above 4 (measured by sun photometer) and an Ångström exponent of -0.08. The observed lidar ratios (particle linear depolarization ratios) in the presented dust cases range from 40.3 sr to 46.9 sr (0.18–0.29) at 355 nm and from 35.7 sr to 42.9 sr (0.31–0.35) at 532 nm wavelength. The particle linear depolarization ratios indicate almost unpolluted dust in the case of a lofted dust layer and pure dust in the near-ground dust cases. The lidar ratio values are lower than typical lidar ratio values for Saharan dust (50–60 sr) and comparable to Middle Eastern/West-Asian dust lidar ratios (35–45 sr). In contrast, the presented case of pollution aerosol of local origin has an Ångström exponent of 2.07 and a lidar ratio (particle linear depolarization ratio) of 55.8 sr (0.03) at 355 nm and 32.8 sr (0.08) at 532 nm wavelength.

## 1 Introduction

Atmospheric mineral dust can be transported over tens of thousands of kilometers away from its sources, arid and semi-arid source regions (Uno et al., 2009; Haarig et al., 2017). More observations, especially in western and Central Asia, are needed to describe global and regional dust transport and to estimate the effect of this dust on air quality (Chin et al., 2007), and climate, via direct and various indirect radiative effects. Extended investigations to Saharan dust close to its source regions (e.g. SAMUM-1,2 (Saharan Mineral Dust Experiment), Fennec climate programm, SHADOW (Study of SaHaran Dust Over West Africa)) (Heintzenberg, 2009; Ansmann et al., 2011a; Ryder et al., 2015; Veselovskii et al., 2016) as well as regarding dust long range transport across the Atlantic ocean (e.g. SALTRACE (Saharan Aerosol Long-range Transport and Aerosol-Cloud-Interaction Experiment)) (Weinzierl et al., 2017) have been conducted to obtain novel data to reduce uncertainties in above mentioned estimations.

However, the global dust belt, which reaches from the Sahara over the Arabian deserts to the Taklamakan and Gobi deserts, contains a lot more arid and semi-arid regions which act as sources for atmospheric mineral dust in the northern hemisphere (Darmenova et al., 2009; Ridley et al., 2016). This dust has a sensitive impact on climate, environmental conditions, ecosystems, and health. Central Asia lies in the middle of this global dust belt and contains some major dust sources (Fig. 1). Therefore, Central Asian countries are frequently affected by atmospheric mineral dust hazardous to respiratory health (Wiggs et al., 2003). Even bacteria, fungi, and viruses can be transported on dust on long distances (Griffin, 2007; Hara and Zhang, 2012; Yamaguchi et al., 2015; Park et al., 2016). Dust is also transported across the highly polluted Central Asia (Balance and Pant, 2003) further eastwards (Tanaka et al., 2005; Mikami et al., 2006) and on its way it is subject to stronger anthropogenic influence compared to the westward transport of Saharan dust (Bi et al., 2016).

Dust from different sources have different mineralogical compositions (Caquineau et al., 2002), different optical properties (Sokolik et al., 1993; Su and Toon, 2011; Schuster et al., 2012; Mamouri et al., 2013), and therefore different impact on radiation and climate (Sokolik and Golitsyn, 1993; Golitsyn et al., 1993b; Nazarov et al., 2010a, b). Central Asia in particular is vulnerable to climate change and is already affected by it (Freedman and Neuzil, 2015; Lioubimtseva and Henebry, 2009; Xu et al., 2016). For example, dramatic glacier shrinking took place in the last decades in Tajikistan (Kayumov, 2010; Chatenoux, 2011), which has an effect on the water resources of Tajikistan and the whole Central Asian area (Chen et al., 2017). The links between mineral dust occurrence and climate change as well as possible feedbacks have to be quantified. Therefore, aerosol profiling in Central Asia is strongly required to better understand regional and global transport and deposition of mineral dust and its effects on radiation budget, cloud and precipitation development, and human health.

The sources and emission of dust in Central Asia, interaction with aerosol pollution, and climate relevant processes are investigated with models and satellite remote sensing (Xi and Sokolik, 2016, 2015a, b; Nobakht et al., 2015). Measurements with sun photometers, despite being rare in inland Central Asia (Klüser et al., 2015; Bayat et al., 2011), can provide column integrated aerosol properties (Semenov et al., 2005; Bi et al., 2016; Chen et al., 2013b), but additionally vertically resolved information on aerosol optical properties are needed to thoroughly investigate radiative effects and dust-cloud-dynamics interaction. The widely used satellite-borne lidar CALIPSO/CALIOP (Cloud-Aerosol Lidar and Infrared Pathfinder Satellite

Observations / Cloud-Aerosol Lidar with Orthogonal Polarization) delivers snapshot-like observations in terms of backscatter profiles at 532 nm and 1064 nm wavelength. To estimate extinction coefficients of dust, the dust lidar ratio is required as an input, which introduces considerable uncertainty (e.g. Wandinger et al., 2010; Tesche et al., 2013; Amiridis et al., 2013). To obtain sophisticated sets of optical properties and detailed information on particle mixture and layering, long-range dust transport, multi-day coherent structures, vertical mixing between the planetary boundary layer (PBL) and lofted aerosol layers, and aerosol-cloud interaction, continuous observations with ground-based Raman lidars or HSRL (High-Spectral-Resolution lidar) are needed. Up to now, only few experiments were performed to characterize the aerosol over Central Asia. Already in 1989 a Soviet-American research team conducted a joint experiment on dust in Tajikistan (Golitsyn and Gillette, 1993; Golitsyn et al., 1993a). Coordinated ground-based, aircraft and satellite measurements took place during two dust storms on 16 and 20 September 1989. The total area occupied by the atmospheric dust during the latter dust storm was approximately  $10^5 \text{ km}^2$  (Smirnov et al., 1993). The mass of dust in the air was estimated as 3.1 t (Smirnov et al., 1994). Chemical analysis of the collected dust (Gomes and Gillette, 1993) showed a low iron content for Central Asian dust. Also, a calcareous character and high contents of soluble salts were observed (Miller-Schulze et al., 2011; Schettler et al., 2014; Groll et al., 2013; Andronova et al., 1993). In general, however, a very high variability (Andronova et al., 1993) lead to considerably different results depending on sampling location. Analysis of quartz filter samples collected in Dushanbe showed, besides significant amounts of dust, strong contribution of anthropogenic black carbon pollution from fuel combustion and smoke (Hansen et al., 1993). Optical and radiative properties have been studied as well (Shukurov et al., 1993). Sokolik and Golitsyn (1993) performed radiative transfer calculations for the dust storms in 1989 and state that the vertical distribution of additional total radiative heating rates can only be calculated with assumptions on vertically homogeneous optical properties and on the height of the dust layer. Sokolik et al. (1993) compared complex refractive indices from the dust collected in Tajikistan with dust from other sources. They found a seemingly different spectral behavior of the complex refractive index of Central Asian dust compared to Saharan dust, but Sokolik et al. (1993) concluded that there are large uncertainties and the discrepancies could be due to different measurement techniques. The measured AOTs during those dust storms were up to 1.3 (Golitsyn and Gillette, 1993) and 3.3 at 550 nm (Pachenko et al., 1993), respectively. A few vertical profiles of the scattering coefficient of the dust were measured with a nephelometer on board an aircraft (Pachenko et al., 1993). These measurements showed that the majority of the dust was concentrated in the lowermost 3–3.5 km below a thermal inversion. Pachenko et al. (1993) mentioned dust plumes reaching through this inversion up to 6 km altitude. The measurements on 16 September 1989 also revealed a dust layer at a height of about 4 km. Fraser (1993) used satellite data to obtain the horizontal distribution of the dust in Central Asia. Fraser (1993) emphasized the need of vertical profiles of the dust mass, which can be estimated from dust extinction coefficient, to determine the movement of the dust. Gillette et al. (1993) modeled dust deposition and transport during this 1989 campaign and mentioned the sensitivity of their model to the dust layer height, especially because of the complicated topography in Tajikistan.

Since recently, a lidar station is operated in Eastern Kyrgyzstan (Chen and Sverdlik, 2007; Chen et al., 2013a). For this station in Kyrgyzstan, Chen et al. (2013b) report dust transported from the Aral Sea region, long-range transported dust from the Middle East as well as dust from the Taklamakan desert. A lidar is operated in Aksu (northwestern Taklamakan) (Kai et al.,

2008), where local dust was measured from ground to a height of 6 km during strong dust outbreaks (Tsunematsu et al., 2005) as well as lofted long-range transported dust layers at heights of 11 km (Mikami et al., 2006). Lidar measurements along the Caspian and Aral Sea and the Lake Balkhash in Kazakhstan were performed by Dieudonné et al. (2015). Dieudonné et al. (2015) showed that dust layers originating in the Caspian and Aral Sea region can regularly spread over wide areas of Russia and last for several days.

Nevertheless, the knowledge of the vertical aerosol distribution over Tajikistan and especially the transport of mineral dust over Central Asia is still insufficient. Therefore, the Central Asian Dust EXperiment (CADEX) was proposed. CADEX provides long-term data on vertical profiles of particle optical properties above Tajikistan. A multiwavelength polarization/Raman lidar (Engelmann et al., 2016) was run continuously in Tajikistan over the 18-months period from March 2015 to August 2016. These measurements are part of the PollyNET (Baars et al., 2016), which is an attempt to contribute to the establishment of a ground-based polarization/Raman lidar network (e.g. Baars, 2011; Kanitz et al., 2013; Giannakaki et al., 2016), with the aim to support regional and global dust modeling.

This article provides an overview of the observations and focuses mainly on four exemplary measurement cases which show typical but different scenarios of atmospheric mineral dust in Tajikistan. The vertical profiles of the measured particle optical properties, calculations of the dust mass concentration, and backward trajectories for the individual cases are presented. In a follow-up publication, a summary of the key findings of the CADEX campaign including a statistical analysis of the entire measurement period will be provided.

In section 2 of this article, the CADEX campaign, the measurement site, the used lidar system, and the used auxiliary data are described. Section 3 gives an overview of the observations and the four exemplary measurement cases are presented in detail. The cases are compared and discussed in section 4.

## 2 The Central Asian Dust EXperiment (CADEX) and instruments

CADEX was a joint German-Tajik project to investigate the atmospheric mineral dust over Tajikistan. The goal of the project was to provide a not yet available data set of optical and microphysical properties of Central Asian mineral dust on a longer-term basis. The data are used to characterize and investigate the particle types, possible aerosol sources, and the aerosol influence on the radiation budget in Central Asia.

Tajikistan lies in the global dust belt in close proximity of some major dust sources like the Taklamakan desert, the Aralkum desert of the desiccating Aral Sea, the Kyzylkum and Karakum deserts east of the Caspian Sea, the Iranian Dasht-e-Kavir and Dasht-e-Lut deserts and the deserts in Afghanistan (Fig. 1). Therefore, Tajikistan is frequently affected by severe dust events and is a net accumulator of dust (Gillette and Dobrowolski, 1993). Tajikistan is a country with a dry continental climate and benefits from its water resources. Melt water discharge from the high mountain regions ensures a high freshwater availability. Furthermore, Central Asia and especially Tajikistan are highly affected by climate change. For example, dramatic glacier shrinking took place in the last decades (Makhmadaliev and Novikov, 2002; Makhmadaliev et al., 2004, 2008; Kayumov and Novikov, 2014). This could have also an effect on the water resources of Tajikistan and the whole Central Asian area in



the future, since the trans-regionally important rivers as Amu Darya and Syr Darya are fed by glacier and snow melt water (Siegfried et al., 2012; Sorg et al., 2014), which are feeding the desiccating Aral Sea, which now became itself a strong dust source (Xi and Sokolik, 2016). On the other hand, deposited dust and anthropogenic black carbon themselves can accelerate glacier melt by altering the glacier's surface albedo (Gabbi et al., 2015; Schmale et al., 2017).

5 The fieldwork of CADEX included column-integrating sun photometer measurements, vertically resolved lidar measurements, and meteorological observations in Dushanbe, Tajikistan. The lidar observations were carried out in Dushanbe at the Polygon measurement site of the Physical Technical Institute of the Academy of Sciences of Tajikistan. The measurement site lies in an urban environment on a hill in the eastern part of Dushanbe. The lidar is installed in front of a laboratory and office building (Fig. 2a). Its coordinates are 38°33'34" N, 68°51'22" E, and its altitude is 864 m ASL. The measurement campaign lasted  
10 from 17 March 2015 until 31 August 2016. During that period at 487 days lidar data has been acquired for a time period of at least 3 h. On 308 of these days the lidar ran even longer than 20 h.

In addition, the field experiment is accompanied by 3-D regional modeling with the regional dust model COSMO-MUSCAT (Wolke et al., 2012; Heinold et al., 2011) and the aerosol-climate model ECHAM-HAM (Zhang et al., 2012; Heinold et al., 2016). The modeling provides a perspective on the sources, transport, as well as the direct radiative effects of measured mineral  
15 dust and associated atmospheric feedbacks. Furthermore, ground-based in situ aerosol measurements were conducted during CADEX to investigate the relationship between the settling of mineral dust along the margin of the Central Asian mountains and the dust transport at high tropospheric levels recorded by lidar. Collection of aerosol (March 2015 to April 2016) was carried out on quartz fiber filters (MK 360, MUNKTELL) using a high-volume filter sampler (DHA-80, DIGITEL) with an inlet for PM<sub>10</sub> (particulate matter with a maximum diameter of 10 µm). Chemical analysis of the collected aerosol was performed  
20 with the techniques described in Fomba et al. (2014). The particle number size distribution was measured with a laser particle counter (GRIMM EDM 180) (Schettler et al., 2014). Dust dry deposition was collected (9 days in August 2016) by means of a flat plate type passive particle collector (Ott and Peters, 2008). Collected particles were subject to electron microscopy single particle analysis (Kandler et al., 2009, 2011), yielding information on particle size distribution, chemical and mineralogical composition as well as mixing state for coarse mode particles with diameter between 0.7 µm and 60 µm. The additional model  
25 and in situ observational data, however, need further analysis and evaluation and will be used in a follow-up study.

## 2.1 Lidar system

The lidar operated in Dushanbe is a multiwavelength polarization/Raman lidar Polly<sup>XT</sup> (PORTabLe Lidar sYstem Polly<sup>XT</sup>, XT stands for extended version, Althausen et al. (2009)) and belongs to PollyNET, a network of permanent or temporary Polly systems (Baars et al., 2016). This specific Polly<sup>XT</sup> was already deployed in numerous field campaigns in the past (e.g. Baars  
30 et al., 2012, 2016). For the CADEX campaign this Polly<sup>XT</sup> was upgraded with a second depolarization channel at 355 nm wavelength (Engelmann et al., 2016). All Polly<sup>XT</sup> systems contain a laser system that emits light at 355 nm, 532 nm, and 1064 nm wavelength by means of frequency doubling and tripling. The receiver of the used system has eight channels and measures the backscattered light at all three emitted wavelengths. The channels at 387 nm, 607 nm, and 407 nm wavelength allow to detect Raman scattering at night time. Another two channels detect cross-polarized light at 355 nm and 532 nm

wavelength. Three times a day, the system performs an automatic  $\Delta 90^\circ$ -depolarization-calibration (Freudenthaler et al., 2009; Freudenthaler, 2016). Figure 2b shows a sketch of the optical layout of the used system. Signals are sampled with a vertical resolution of 7.5 m and are stored with a temporal resolution of 30 s. Full overlap is reached at about 1.5 km AGL (Baars et al., 2012). The resulting products of the Polly<sup>XT</sup> are summarized in Table 1. The corresponding uncertainties, discussed in detail in Freudenthaler et al. (2009), Baars et al. (2012, 2016) and Engelmann et al. (2016), result from uncertain input parameters and signal noise and are given as error bars in the results section. The polarization lidar photometer networking method (POLIPHON, (Ansmann et al., 2011b, 2012)) is applied to separate dust and non-dust optical and related microphysical properties. This separation is performed with respect to the particle backscatter coefficient (Teschke et al., 2009a; Groß et al., 2011; Mamouri and Ansmann, 2014) using the measured total particle linear depolarization ratio at 532 nm wavelength as well as typical particle linear depolarization ratio values (Teschke et al., 2009a; Müller et al., 2007) for dust (0.31) and non-dust (0.05). For the conversion of particle extinction coefficients into particle volume concentration, conversion factors for dust ( $0.61 \cdot 10^{-6}$  m) and non-dust ( $0.41 \cdot 10^{-6}$  m) from extended AERONET data analysis from Mamouri and Ansmann (2017) are used. Together with the densities of dust ( $2.6 \text{ g cm}^{-3}$ ) and non-dust ( $1.6 \text{ g cm}^{-3}$ ), this allows to calculate dust mass concentration profiles using characteristic lidar ratio values of dust (40 sr) and non-dust (80 sr). The uncertainty of such an estimation is given as 40% (Ansmann et al., 2011b). Bravo-Aranda et al. (2015) compared the POLIPHON method directly to in situ aircraft measurements and found less than 30% relative difference. The measured lidar data is uploaded daily and processed automatically (Baars et al., 2016) to be displayed as quicklooks on the PollyNET homepage (<http://polly.rsd.tropos.de/>).

## 2.2 Auxiliary data

Co-located with the lidar, a sun photometer as part of the AErosol RObotic NETwork (AERONET) (Holben et al., 1998) was operated. The Dushanbe AERONET station ([http://aeronet.gsfc.nasa.gov/cgi-bin/type\\_one\\_station\\_opera\\_v2\\_new?site=Dushanbe](http://aeronet.gsfc.nasa.gov/cgi-bin/type_one_station_opera_v2_new?site=Dushanbe)) is operational since 2010 (Abdullaev et al., 2012). The CIMEL sun photometer measures at eight wavelengths (1020, 870, 675, 500, 490, 440, 380, and 340 nm) and retrieves the AOT and further column integrated particle optical and microphysical properties.

To calculate the Rayleigh contribution to the lidar signals in order to obtain particle optical properties, profiles of temperature and pressure are needed. Profiles of the GDAS (Global Data Assimilation System) with  $1^\circ$  spatial resolution from the National Weather Service's National Centers for Environmental Prediction (NCEP) at the coordinates  $39^\circ$  N,  $69^\circ$  E were used (GDAS, 2017), because no radiosonde launches were performed in Tajikistan since several years.

Publicly available trajectory models were used to assess the possible source regions and transport routes of the dust. The HYSPLIT model (Hybrid Single Particle Lagrangian Integrated Trajectory Model) (Stein et al., 2015; Rolph, 2016) as well as the FLEXPART model (FLEXible PARTicle dispersion model) version 9 (Stohl et al., 2005) were run to calculate backward trajectories for the presented example cases. The  $0.5^\circ$  GDAS meteorology for HYSPLIT and the  $1^\circ$  GFS (Global Forecast System) meteorology for FLEXPART is used. The HYSPLIT and FLEXPART backward trajectories were calculated with a starting time approximately corresponding to the time of observation of the aerosol layer by lidar. The arrival heights above the measurement site were set to the base, center, and top of the observed layer. The FLEXPART backward trajectories were

calculated for the arrival height above the measurement site in the height range of the aerosol layer measured with the lidar. The model output is the common logarithm of the accumulated residence time of air masses below 3 km during the model run time (Schwarz, 2015). This 3 km height is chosen because air is likely loaded with dust in the source regions, where planetary boundary layers of this magnitude occur. The spatial resolution of the model output is  $1^\circ$ .

## 5 3 Observations

### 3.1 Overview

An overview of the daily means of the AOT at 500 nm wavelength and the Ångström exponent from the 440–870 nm spectral range is shown in Fig. 3. The data were measured by sun photometer from March 2015 until August 2016 (cloudscreened AERONET data, Level 1.5). For the six 3-months seasons of the measurement campaign the means of the AOT, the Ångström exponent and the fine mode fraction from the 500 nm measurements are shown in Tab. 2. A seasonal transition is obvious from the winter to the summer months. The AOT increases from spring to summer and decreases again in autumn. The Ångström exponent and the fine mode fraction behave reversely. There is a dusty 6-months season from April to September, and a less dusty season from October-March, but a strong dust event still occurred in October 2015. Within these seasons strong variations in the AOT and especially in the Ångström exponent occur from day to day. This indicates fast changes from dust dominated to pollution dominated aerosol conditions.

Gkikas et al. (2009, 2013, 2016) and Georgoulas et al. (2016) classified dust events in the Mediterranean region based on the long-term mean of the AOT measured by satellite and ground stations. According to this classification, a strong dust event equals or exceeds the mean AOT plus 2 times the standard deviation, while an extreme dust event equals or exceeds the mean AOT plus 4 times the standard deviation. This classification can be pursued for the station in Dushanbe with AERONET data (Level 2.0), which are available from July 2010 to November 2016 (1413 days). The mean of the daily mean AOT at 500 nm wavelength is 0.28 with a standard deviation of 0.23. As a comparison, Georgoulas et al. (2016) found a long-term mean AOT at 550 nm wavelength of  $0.22 \pm 0.19$  for the Eastern Mediterranean region. So in Dushanbe, days with an  $\text{AOT} \geq 0.74$  are therefore strong dust events, while days with an  $\text{AOT} \geq 1.2$  at 500 nm wavelength are extreme dust events. Considering only at the measurement period from March 2015 to August 2016, the mean AOT at 500 nm wavelength was 0.28 (with a standard deviation of 0.16) and therefore equal to the long-term mean. According to this classification, 6 episodes of strong dust and 6 episodes of extreme dust occurred during the measurement period. In the period July 2010 to November 2016 (6.25 years) strong dust events occurred 19 times (3 per year), extreme dust events 10 times (1.6 per year) and more than extreme dust events with an AOT above 2 at 500 nm wavelength occurred 4 times (0.6 per year).

### 3.2 Illustrative measurement examples

To illustrate the frequently observed variety and complexity of dust layers above Tajikistan, examples of dust layers of various origin occurring at all heights from the surface to cirrus level are presented in Fig. 4. Figure 4a shows the temporal development

of the range-corrected signal of the 1064 nm wavelength channel from 20–24 April 2015. During these four days multiple dust layers arrived above Tajikistan. On 20 April 2015 there were aerosol layers up to 4 km height. At the same time, descending dust layers crossed the lidar station at heights of about 10 km. On 22 April 2015 a second dust layer arrived at 10 km height, which again descended. According to FLEXPART trajectories, the Arabian Peninsula, Iran, and also parts of Uzbekistan and Turkmenistan were source regions for the detected dust (Fig. 4d). A high altitude aerosol layer was measured on 13 May 2015, reaching roughly from 6.5 km to 10 km altitude (Fig. 4b). FLEXPART backtrajectory analysis shows that this high layer was long-range transported towards the measurement site from the Mediterranean/North African region (Fig. 4e). The third scenario is a near-ground dust layer (Fig. 4c, described in detail in Sec. 3.3.2), which was measured on 8 and 9 August 2015 and had Central Asian sources (Fig. 4f).

### 10 3.3 Case studies

In the following sections we discuss four strongly contrasting cases in more detail: (1) A lofted layer of Middle Eastern dust that occurred on 13 April 2015 (Case 1, Fig. 5, section 3.3.1), (2) an extreme dust event with Central Asian dust, which was recorded on 8 and 9 August 2015 (Case 2, Fig. 6, section 3.3.2), (3) the most extreme dust event during the CADEX campaign with dust from Central Asian sources, which was observed on 14 July 2016 (Case 3, Fig. 7, section 3.3.3), and (4) a contrasting case with a pollution layer of local origin, which was measured on 13 May 2015 (Case 4, Fig. 8, section 3.3.4).

#### 3.3.1 Case 1: 13 April 2015, lofted dust layer

Figure 5a shows the temporal development of the range-corrected signal of the 1064 nm wavelength channel on 12–13 April 2015. The dust layer arrived over Dushanbe on 12 April 2015 at an altitude between 5 km and 6 km. The slowly descending dust layer (red) contained dense clouds (grey). On 13 April 2015 the dust layer extended at about 2.5–5 km height. Its base was very sharp while on top of the layer a thinner aerosol layer was measured up to 7 km. At the end of that day, a down-mixing of dust into the lowermost 2 km occurred. The HYSPLIT backward trajectories arriving on 13 April 2015 show that the air parcels are coming from southwestern direction (Fig. 5d). The upper and the lower trajectories come from the Arabian Peninsula and travel over central Iran and along the Afghanistan-Turkmenistan border. The upper trajectory is always more than 3 km above ground, while the lower trajectory reached below 2 km once above the Arabian Peninsula. The center trajectory is coming from more southern direction, traveling over southern Iran, where it reaches lowest down to below 1 km above the Iranian Dashte-Lut desert.

Similar to HYSPLIT, the FLEXPART backward trajectories indicate Iran including its southern parts and large parts of the Arabian Peninsula as source regions (Fig. 5c). In contrast to HYSPLIT, the FLEXPART backward trajectories resided partly above Afghanistan.

Vertical profiles of the optical properties of this dust layer are shown in Fig. 5b. The particle linear depolarization ratio at 532 nm wavelength is on average 0.31 and 0.34 at its maximum. This shows that this lofted dust layer consists of almost unpolluted dust. Depolarization ratios of pure Middle Eastern dust close to its source regions have been found as 0.3–0.35 at 532 nm wavelength (Mamouri et al., 2013, 2016). The mean lidar ratio is  $35.7 \pm 1.7$  sr at 532 nm wavelength, what is

comparable to measurements for Middle Eastern dust (Mamouri et al., 2013; Nisantzi et al., 2015). The lofted dust layer has a maximum dust mass concentration of  $196 \mu\text{g m}^{-3}$  at the bottom of the layer at 2.7 km. In the middle of the layer at 3.5 km the dust mass concentration reaches  $171 \mu\text{g m}^{-3}$ . The integrated dust mass up to 5 km altitude is  $0.51 \text{ g m}^{-2}$ . The non-dust components in this layer are negligible, except in the upper parts of the layer, where the particle linear depolarization ratio partly drops below 0.31 at 532 nm wavelength.

### 3.3.2 Case 2: 8/9 August 2015, extreme dust event

An extreme dust event was measured on 8 and 9 August 2015. Figure 4c shows the temporal development of the range-corrected signal of the 1064 nm wavelength channel on 8–9 August 2015. The dust got lifted up with the evolution of the convective boundary layer. This lead to a very strong signal in the lowest 1.5 km with dust extending up to about 3 km height.

The FLEXPART trajectories arriving on 8 August 2015 (Fig. 4f) exhibit a large area northwest of Tajikistan with high accumulated residence times, even west of the Caspian Sea.

Vertical profiles of the optical properties of this dust layer are shown in Fig. 6. The particle extinction coefficients at 355 nm and 532 nm wavelength are about  $550 \text{ Mm}^{-1}$  at 1 km height. Apparently, the layer is divided in two parts. There is a maximum in the first 1 km above ground and then again at about 2.5 km height. But looking at the intensive properties, especially the particle linear depolarization ratio and the extinction-related Ångström exponent, continuous features are recognizable. The twofold layer probably contains the same aerosol and originates from the same source region since it has almost identical optical properties. The apparent layer boundary at 1.6 km is probably formed by the diurnal cycle of the boundary layer. The development of the convective boundary layer after about 06:00 UTC on 8 August 2015 is visible in Fig. 4c. The mean particle linear depolarization ratios are higher than in the lofted dust layer before (Case 1), on average 0.35 at 532 nm wavelength. The maximum dust mass concentration of that extreme near-ground dust layer is  $845 \mu\text{g m}^{-3}$  at 0.65 km height, in the middle of the layer the minimum is  $475 \mu\text{g m}^{-3}$  at 1.7 km. In the upper part at 2.4 km the dust mass concentration is again  $663 \mu\text{g m}^{-3}$ . The integrated dust mass up to 4 km altitude is  $3.2 \text{ g m}^{-2}$ .

### 3.3.3 Case 3: 14 July 2016, most extreme dust event

The most extreme dust event during the CADEX campaign occurred on 14 July 2016. The dust persisted during the next four days (Fig. 3). The FLEXPART trajectories arriving on 14 July 2016 (Fig. 7b) show large accumulated residence times as far away as western Iran, but the highest values are in Uzbekistan. The accumulated residence times above Kazakhstan are high, but they are also reaching eastwards towards the Lake Balkhash, differently than on 8 August 2015 (Fig. 4f). The HYSPLIT backward trajectories arriving on 14 July 2016 are shown in Fig. 7c. The higher trajectory arrived from the Aral Sea and the lower one from central Kazakhstan through Turkmenistan and Afghanistan. The vertical profiles of the measured optical properties of this dust layer are presented in Fig. 7a. The particle extinction coefficients at 355 nm and 532 nm wavelength are about  $1.7 \text{ km}^{-1}$  at 1 km height. The resulting AOT is 3.89 at 532 nm wavelength, what is much higher than the long-term mean plus 10 times the standard deviation (see section 3.1). The 14 July 2016 has the highest daily mean AOT since the beginning of the record in Dushanbe with the AERONET sun photometer in 2010. So this dust event could well be called record-breaking

like Mamouri et al. (2016) called an extreme dust event in September 2015 in the Mediterranean. The mean particle linear depolarization ratio at 355 nm wavelength is 0.29 and thus higher than in Case 2, while at 532 nm wavelength it is equally 0.35. The calculated dust mass concentrations for that most extreme dust is largest at 1 km altitude with  $2.8 \text{ mg m}^{-3}$ . The integrated dust mass up to 2.7 km altitude is  $6.5 \text{ g m}^{-2}$ .

### 5 3.3.4 Case 4: 13 May 2015, contrasting case with local pollution

Figure 4b shows the temporal development of the range-corrected signal of the 1064 nm wavelength channel on 13–14 May 2015. There were several aerosol layers distributed up to 10 km height. The HYSPLIT backward trajectories arriving on 13 May 2015 are shown in Fig. 8b. The upper trajectory is coming from northwest of Tajikistan and the lower trajectories reach down to the ground the same day inside Tajikistan, indicating the contribution of local pollution to the lower altitude aerosol

10 layers.

The vertical profiles of the measured optical properties are presented in Fig. 8a. The particle extinction coefficients in the layer at about 2–4 km are about  $40 \text{ Mm}^{-1}$  at 355 nm and  $20 \text{ Mm}^{-1}$  at 532 nm wavelength. This aerosol layer is barely depolarizing and the mean extinction-related Ångström exponent is  $2.07 \pm 0.72$ . The dust mass concentration reaches  $6.9 \mu\text{g m}^{-3}$  at 2.5 km height, the non-dust mass concentration is at the same time dominating with  $30.7 \mu\text{g m}^{-3}$ . The integrated dust mass up to 15 10.4 km altitude is  $0.05 \text{ g m}^{-2}$  and the integrated non-dust mass is  $0.13 \text{ g m}^{-2}$ . Finally, an overview over the individual cases is given in Tab. 3, where the intensive optical properties and AOTs of the presented example cases are summarized.

## 4 Discussion

Four exemplary measurement cases are described in this article. Two of them are from spring and one is from summer 2015 and one is from summer 2016. In early spring, the AOT was predominantly low ( $\leq 0.3$ ) (Fig. 3) but a lot of dust layers occurred 20 (e.g. 13 April 2015 (Case 1, Fig. 5), 20–24 April 2015 (Fig. 4a). In summer, the AOT in general was higher ( $\geq 0.35$ ) and some dust events overtopped that clearly (e.g. 8–9 August 2015 (Case 2, Fig. 6), 14 July 2016 (Case 3, Fig. 7)). However, a statistical analysis of the lidar profiles of the whole measurement period has yet to follow.

### 4.1 Dust AOT and mass concentration

The lofted dust layer in Case 1 contributed significantly to the total AOT of about 0.4 at 532 nm wavelength. Case 2 had an AOT 25 of 1.5 at 532 nm wavelength. The most extreme dust case during the measurement campaign was Case 3 with an extraordinary AOT of above 3.5 at 532 nm wavelength measured by lidar, respectively above 4 by sun photometer. The AOT of Case 2 is comparable to a dust event during the Soviet-American campaign in 1989 (AOT of 1.3 at 550 nm wavelength) (Gomes and Gillette, 1993). The most extreme dust event (Case 3) is comparable to the second dust event during the Soviet-American campaign (AOT of 3.3 at 550 nm wavelength) (Pachenko et al., 1993). Pachenko et al. (1993) estimated the maximum AOT of 30 that dust event a day before was greater than 10, based on horizontal visibility, which was reported to be 50–200 m (Smirnov et al., 1994). As a comparison, Mamouri et al. (2016) reported a visibility of 500–600 m for an extraordinary dust event in the

Mediterranean area, inferring an AOT of 4.8 to 9 depending on the layer height. This again lead to an estimate of the column dust load of 8–15 g m<sup>-2</sup>. These values are higher than for Case 3, where the AOT was 3.89 at 532 nm wavelength, and the integrated dust mass was 6.7 g m<sup>-2</sup>. The visibility calculated from the measured particle extinction coefficient of 1.7 km<sup>-1</sup> (measured at 0.87 km altitude) was 2.3 km. The AOT and the dust mass concentrations for this Case 3 were calculated based on the lidar data only, but the sun photometer measured even a higher AOT of 4.45 at 500 nm wavelength.

## 4.2 Dust layer height

The multiple high altitude dust layers preceding the lofted dust layer on 24 April 2015 (Fig. 4a) as well as the high aerosol layer on 13 May 2015 (Fig. 4b) reached heights of about 10 km AGL. This is higher than the highest point of the Pamir mountains (7.6 km ASL), which means that these layers can cross the Pamir or the Tianshan mountains (7.4 km ASL) and can be transported further eastwards. A similar observation was made in spring 2003 when layers of elevated depolarization at 9–11 km height have been measured by lidar in Aksu (northwestern Taklamakan) (Mikami et al., 2006). Within a week time, lidar stations in Japan measured dust at altitudes between 2 km and 6 km, without having major dust outbreaks in the Taklamakan and Gobi deserts. Model simulations indicated that the dust was transported to Japan via north of the Tianshan mountains (Tanaka et al., 2005). For that case, Tanaka et al. (2005) estimated 50% of the dust particles arriving in Japan came from the Sahara, 30% from the Middle East, and only 10% from China.

## 4.3 Lidar ratio and dust source region

The observed lidar ratios in the presented dust cases (Cases 1,2,3) range from 40.3 sr to 46.9 sr at 355 nm and from 35.7 sr to 42.9 sr at 532 nm wavelength. The lofted dust layer (Case 1) has lidar ratios of 42.2 sr at 355 nm and 35.7 sr at 532 nm wavelength. The near-ground dust layers (Cases 2,3) have lidar ratios of 46.9 sr and 40.3 sr at 355 nm and 42.9 sr and 38.7 sr at 532 nm wavelength. The only weakly depolarizing pollution aerosol in Case 4 has lidar ratios of 55.8 sr at 355 nm and 32.8 sr at 532 nm wavelength, but with high variability.

Saharan dust is found to have lidar ratios of 50–60 sr at 532 nm wavelength (e.g. Groß et al., 2015; Tesche et al., 2009b). Schuster et al. (2012) used AERONET data sets to retrieve pure dust lidar ratios and found lower values for West-Asian dust than for North African dust. Similar results were found from Raman lidar measurements in Cyprus with lidar ratios of 35–45 sr for Middle Eastern dust (Mamouri et al., 2013; Nisantzi et al., 2015). In east Asia, lidar ratios of Asian dust of for example 47 ± 18 sr (Sakai et al., 2003) and 42–73 sr (Liu et al., 2002) at 532 nm wavelength were reported. In China, lidar ratios of 40 ± 5 sr (Tesche et al., 2007) and 35 ± 5 sr (Müller et al., 2007) at 532 nm wavelength for dust from the Gobi desert were measured. The situation in Central Asia is even more unclear, as almost no measurements exist. Although direct measurements of dust lidar ratios inside of the Taklakaman are not yet available, Jin et al. (2010) used constrains to retrieve a dust lidar ratio of 42 ± 3 sr at 532 nm wavelength in Aksu (northwestern Taklamakan). Chen et al. (2013b) measured very low lidar ratios of 8–29 sr at 532 nm wavelength in weakly depolarizing dust layers in Kyrgyzstan. For dust from the Caspian and Aral sea region, a lidar ratio of 43 ± 3 sr at 532 nm was measured (Dieudonné et al., 2015).

The lidar ratios of the example dust cases (Cases 1,2,3) agree well with these values. There is no clear difference between the

lofted dust layer (Case 1), which is long-range transported Middle Eastern dust and the near-ground fresh dust layers (Case 2,3) from Central Asian sources. The lofted dust layer in Case 1 has a very low lidar ratio of  $35.7 \pm 1.7$  at 532 nm wavelength, indeed comparable to the ones ( $33.7 \pm 6.7$  to  $39.1 \pm 5.1$  sr at 532 nm) found by Mamouri et al. (2013) for dust from the Middle East. On the other hand, also the most extreme dust event (Case 3) has a lidar ratio below 40 sr at 532 nm wavelength.

5 Then again, slightly larger lidar ratios above 40 sr were measured in the extreme near-ground dust layer (Case 2). The lidar ratios of the pollution layer (Case 4) differ significantly between the two wavelengths.

#### 4.4 Dust particle linear depolarization ratio

The observed depolarization ratios in the presented dust cases (Cases 1,2,3) range from 0.18 to 0.29 at 355 nm and from 0.31 to 0.35 at 532 nm wavelength. The lofted dust layer (Case 1) has a depolarization ratio of 0.18 at 355 nm and 0.31 at 532 nm wavelength. The two near-ground dust cases (Cases 2,3) have depolarization ratios of 0.23 and 0.29 at 355 nm and both 0.35 at 532 nm wavelength. In contrast to the dust cases, the local pollution aerosol (Case 4) has a low depolarization ratio of 0.03 at 355 nm and 0.08 at 532 nm wavelength.

Dieudonné et al. (2015) measured a particle linear depolarization ratio of  $0.23 \pm 0.02$  at 532 nm wavelength for lofted dust from the Caspian and Aral sea region. From the Kyrgyzstan station depolarization ratio values of 0.1–0.15 for lofted dust layers (Chen et al., 2013b) and around 0.2 at 532 nm wavelength for near-ground dust layers were reported (Chen and Sverdlík, 2007). But it has to be considered that this station is already located at 1.9 km ASL. In Aksu (northwestern Taklamakan), Kai et al. (2008) measured depolarization ratios at 532 nm wavelength of  $0.09^*–0.11^*$  in a lofted dust layer and  $0.18^*–0.33^*$  in a near-ground dust layer. Iwasaka et al. (2003) measured a depolarization ratio of  $0.27^*$  at 532 nm wavelength in a lofted dust layer in Dunhuang (northern Taklamakan). Note that the values denoted with \* are published as aerosol depolarization potentials and are converted to particle linear depolarization ratios (see Burton et al., 2014; Cairo et al., 1999; Gimmestad, 2008).

The measured depolarization ratios for the presented dust cases (Cases 1,2,3) are mostly higher than these literature values for dust measured in or close to Central Asia. This suggests that those studies might have described also observations of polluted/mixed dust. The range of the presented particle depolarization ratios (0.31–0.35) at 532 nm wavelength is comparable with the values of fresh Saharan dust (0.27–0.35) (Freudenthaler et al., 2009) or Middle Eastern dust (0.25–0.32) (Mamouri et al., 2016). The spectral difference between the wavelengths is considerable. The range of the presented particle depolarization ratios at 355 nm wavelength (0.18–0.29) is large, with the exceptionally high value during the most extreme dust event (Case 3). In Saharan dust, particle depolarization ratios at 355 nm wavelength in ranges of 0.22–0.31 were measured at the source region (Freudenthaler et al., 2009) and 0.21–0.27 after long-range transport (Burton et al., 2015; Groß et al., 2015; Haarig et al., 2016, 2017). The near-ground dust cases (Cases 2,3) had higher particle depolarization ratios at both wavelengths than the lofted dust layer of Case 1. Nevertheless, the lofted dust layer in Case 1 had a depolarization ratio of 0.31 at 532 nm wavelength, which indicates only minor changes of Middle Eastern dust depolarization characteristics during its long-range transport towards Central Asia. On the other hand, the lofted dust layer in Case 1 had a lower depolarization ratio at 355 nm wavelength than long-range transported Saharan dust.



## 5 Conclusions

For the first time, multiwavelength polarization/Raman lidar observations have been carried out in Tajikistan in the framework of the CADEX campaign. The continuous 18-months measurement provide a unique data set of vertically resolved aerosol optical properties in Central Asia. Additional ground-based in situ measurements and modeling studies accompanied the CADEX campaign (not presented here). During the campaign, dust layers were observed frequently from the surface to tropopause heights, having source regions in Africa, Middle East and Central Asia. Four case studies which are typical for the conditions at the location have been presented, such as lofted dust layers and extreme dust events. As a contrast to the dust cases, an example of a pollution aerosol layer of local origin illustrates the non-negligible anthropogenic influence on the aerosol in Tajikistan. The observed particle linear depolarization ratios for the dust cases range from 0.18–0.29 at 355 nm and 0.31–0.35 at 532 nm wavelength. The presented examples of near-ground layers of Central Asian dust had high depolarization ratios in both wavelengths indicating pure dust conditions. The presented lofted dust layer of long-range transported Middle Eastern dust had slightly lower depolarization ratios, especially at 355 nm wavelength. Nevertheless, the values still indicate almost unpolluted dust conditions after long-range transport to Central Asia. The observed lidar ratios in the presented dust cases range from 40.3 sr to 46.9 sr at 355 nm and from 35.7 sr to 42.9 sr at 532 nm wavelength. These lidar ratio values are lower than typical lidar ratios of Saharan dust (50–60 sr) and comparable to lidar ratios of dust from Middle East/West-Asia (35–45 sr). Further analyses of the data set, to be published in a follow-up publication, include a statistical analysis of the whole measurement period regarding the optical properties, dust mass concentrations, dust layer heights, seasons, and source regions.

Nevertheless, more measurements in Central Asia are needed to analyze long-term trends, especially with respect to climate change. Such measurements offer also new satellite comparison possibilities. Moreover, the assimilation of these data into dust models (regional and global transport, aerosol optical properties, radiative transfer) will help to reduce uncertainties. Therefore, a permanent lidar station in Tajikistan will be established starting most likely from 2019.

## 6 Data availability

HYSPLIT backward trajectories are calculated via the available online tools (<http://ready.arl.noaa.gov/HYSPLIT.php>). AERONET sun photometer data are available from the AERONET web page ([http://aeronet.gsfc.nasa.gov/cgi-bin/type\\_one\\_station\\_opera\\_v2\\_new?site=Dushanbe](http://aeronet.gsfc.nasa.gov/cgi-bin/type_one_station_opera_v2_new?site=Dushanbe)). Data of the FLEXPART backward trajectory calculations and the CADEX Polly<sup>XT</sup> lidar data are available at Leibniz Institute for Tropospheric Research, Leipzig.

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## References

- Abdullaev, S. F., Nazarov, B. I., Salikhov, T., , and Maslov, V. A.: Correlations of surface air temperature and optical thickness of arid aerosol according to AERONET data, *Atmos. Ocean. Opt.*, 25, 428–433, 2012.
- Althausen, D., Engelmann, R., Baars, H., Heese, B., Ansmann, A., Müller, D., and Komppula, M.: Portable Raman Lidar Polly<sup>XT</sup> for Automated Profiling of Aerosol Backscatter, Extinction, and Depolarization, *J. Atmos. Oceanic Tech.*, 26, 2366–2378, doi:10.1175/2009JTECHA1304.1, 2009.
- Amiridis, V., Wandinger, U., Marinou, E., Giannakaki, E., Tsekeri, A., Basart, S., Kazadzis, S., Gkikas, A., Taylor, M., Baldasano, J., and Ansmann, A.: Optimizing CALIPSO Saharan dust retrievals, *Atmos. Chem. Phys.*, 13, 12 089–12 106, doi:10.5194/acp-13-12089-2013, 2013.
- Andronova, A. V., Gomes, L., Smirnov, V. V., Ivanov, A. V., and Shukurova, L. M.: Physico-chemical characteristics of dust aerosols deposited during the Soviet-American experiment (Tadzhikistan, 1989), *Atmos. Environ.*, 27, 2487–2493, doi:10.1016/0960-1686(93)90020-Y, 1993.
- Ansmann, A., Petzold, A., Kandler, K., Tegen, I., Wendisch, M., Müller, D., Weinzierl, B., Müller, T., and Heintzenberg, J.: Saharan Mineral Dust Experiments SAMUM–1 and SAMUM–2: what have we learned?, *Tellus B*, 63, 403–429, doi:10.1111/j.1600-0889.2011.00555.x, 2011a.
- Ansmann, A., Tesche, M., Seifert, P., Groß, S., Freudenthaler, V., Apituley, A., Wilson, K. M., Serikov, I., Linné, H., Heinold, B., Hiebsch, A., Schnell, F., Schmidt, J., Mattis, I., Wandinger, U., and Wiegner, M.: Ash and fine-mode particle mass profiles from EARLINET-AERONET observations over central Europe after the eruptions of the Eyjafjallajökull volcano in 2010, *J. Geophys. Res. Atmos.*, 116, doi:10.1029/2010JD015567, 2011b.
- Ansmann, A., Seifert, P., Tesche, M., and Wandinger, U.: Profiling of fine and coarse particle mass: case studies of Saharan dust and Eyjafjallajökull/Grimsvötn volcanic plumes, *Atmos. Chem. Phys.*, 12, 9399–9415, doi:10.5194/acp-12-9399-2012, 2012.
- Baars, H.: Aerosol profiling with lidar in the Amazon Basin during the wet and dry season 2008, Ph.D. thesis, University of Leipzig, 2011.
- Baars, H., Ansmann, A., Althausen, D., Engelmann, R., Heese, B., Müller, D., Artaxo, P., Paixao, M., Pauliquevis, T., and Souza, R.: Aerosol profiling with lidar in the Amazon Basin during the wet and dry season, *J. Geophys. Res. Atmos.*, 117, doi:10.1029/2012JD018338, 2012.
- Baars, H., Kanitz, T., Engelmann, R., Althausen, D., Heese, B., Komppula, M., Preißler, J., Tesche, M., Ansmann, A., Wandinger, U., Lim, J.-H., Ahn, J. Y., Stachlewska, I. S., Amiridis, V., Marinou, E., Seifert, P., Hofer, J., Skupin, A., Schneider, F., Bohlmann, S., Foth, A., Bley, S., Pfüller, A., Giannakaki, E., Lihavainen, H., Viisanen, Y., Hooda, R. K., Pereira, S. N., Bortoli, D., Wagner, F., Mattis, I., Janicka, L., Markowicz, K. M., Achtert, P., Artaxo, P., Pauliquevis, T., Souza, R. A. F., Sharma, V. P., van Zyl, P. G., Beukes, J. P., Sun, J., Rohwer, E. G., Deng, R., Mamouri, R. E., and Zamorano, F.: An overview of the first decade of Polly<sup>NET</sup>: an emerging network of automated Raman-polarization lidars for continuous aerosol profiling, *Atmos. Chem. Phys.*, 16, 5111–5137, doi:10.5194/acp-16-5111-2016, 2016.
- Balance, R. and Pant, B. D., eds.: Environment Statistic in Central Asia: Progress and Prospects, ERD Working Paper Series No. 36, Asian Development Bank Manila, Philippines, 2003.
- Bayat, A., Masoumi, A., and Khalesifard, H. R.: Retrieval of atmospheric optical parameters from ground-based sun-photometer measurements for Zanjan, Iran, *Atmos. Meas. Tech.*, 4, 857–863, doi:10.5194/amt-4-857-2011, 2011.
- Bi, J., Huang, J., Holben, B., and Zhang, G.: Comparison of key absorption and optical properties between pure and transported anthropogenic dust over East and Central Asia, *Atmos. Chem. Phys.*, 16, 15 501–15 516, doi:10.5194/acp-16-15501-2016, 2016.

- Bravo-Aranda, J. A., Titos, G., Granados-Muñoz, M. J., Guerrero-Rascado, J. L., Navas-Guzmán, F., Valenzuela, A., Lyamani, H., Olmo, F. J., Andrey, J., and Alados-Arboledas, L.: Study of mineral dust entrainment in the planetary boundary layer by lidar depolarisation technique, *Tellus B*, 67, 26 180, doi:10.3402/tellusb.v67.26180, 2015.
- Burton, S. P., Vaughan, M. A., Ferrare, R. A., and Hostetler, C. A.: Separating mixtures of aerosol types in airborne High Spectral Resolution Lidar data, *Atmos. Meas. Tech.*, 7, 419–436, doi:10.5194/amt-7-419-2014, 2014.
- Burton, S. P., Hair, J. W., Kahnert, M., Ferrare, R. A., Hostetler, C. A., Cook, A. L., Harper, D. B., Berkoff, T. A., Seaman, S. T., Collins, J. E., Fenn, M. A., and Rogers, R. R.: Observations of the spectral dependence of linear particle depolarization ratio of aerosols using NASA Langley airborne High Spectral Resolution Lidar, *Atmos. Chem. Phys.*, 15, 13 453–13 473, doi:10.5194/acp-15-13453-2015, 2015.
- Cairo, F., Di Donfrancesco, G., Adriani, A., Pulvirenti, L., and Fierli, F.: Comparison of various linear depolarization parameters measured by lidar, *Appl. Opt.*, 38, 4425–4432, doi:10.1364/AO.38.004425, 1999.
- Caquineau, S., Gaudichet, A., Gomes, L., and Legrand, M.: Mineralogy of Saharan dust transported over Northwestern tropical Atlantic Ocean in relation to source regions, *J. Geophys. Res. Atmos.*, 107, doi:10.1029/2000JD000247, 2002.
- Chatenoux, B.: Central Asia Glaciers' Study - Current state of knowledge and recommendations, Central Asian Climate Risk Management Programme (CA-CRM), Literature and data review, gaps' identification and recommendations for future work prepared by UNEP/DEWA/GRID-Europe, 2011.
- Chen, B. and Sverdlik, L.: Optical and microphysical characteristics of aerosol structures in Central Asia, in: International Conf. on Lasers, Applications, and Technologies 2007: Environmental Monitoring and Ecological Applications; Optical Sensors in Bio, Chemical, and Engineering Technologies; and Femtosecond Laser Pulse Filamentation, vol. 6733, doi:10.1117/12.753117, 2007.
- Chen, B. B., Sverdlik, L. G., Imashev, S. A., Solomon, P. A., Lantz, J., Schauer, J. J., Shafer, M. M., Artamonova, M. S., and Carmichael, G.: Lidar Measurements of the Vertical Distribution of Aerosol Optical and Physical Properties over Central Asia, *Air Qual. Atmos. Health*, p. 385–396, doi:10.1007/s11869-012-0192-5, 2013a.
- Chen, B. B., Sverdlik, L. G., Imashev, S. A., Solomon, P. A., Lantz, J., Schauer, J. J., Shafer, M. M., Artamonova, M. S., and Carmichael, G. R.: Lidar Measurements of the Vertical Distribution of Aerosol Optical and Physical Properties over Central Asia, *Int. J. Atmos. Sci.*, doi:10.1155/2013/261546, 2013b.
- Chen, Y., Li, W., Fang, G., and Li, Z.: Review article: Hydrological modeling in glacierized catchments of central Asia – status and challenges, *Hydrol. Earth. Syst. Sc.*, 21, 669–684, doi:10.5194/hess-21-669-2017, 2017.
- Chin, M., Diehl, T., Ginoux, P., and Malm, W.: Intercontinental transport of pollution and dust aerosols: implications for regional air quality, *Atmos. Chem. Phys.*, 7, 5501–5517, doi:10.5194/acp-7-5501-2007, 2007.
- Darmenova, K., Sokolik, I. N., Shao, Y., Marticorena, B., and Bergametti, G.: Development of a physically based dust emission module within the Weather Research and Forecasting (WRF) model: Assessment of dust emission parameterizations and input parameters for source regions in Central and East Asia, *J. Geophys. Res. Atmos.*, 114, doi:10.1029/2008JD011236, 2009.
- Dieudonné, E., Chazette, P., Marnas, F., Totems, J., and Shang, X.: Lidar profiling of aerosol optical properties from Paris to Lake Baikal (Siberia), *Atmos. Chem. Phys.*, 15, 5007–5026, doi:10.5194/acp-15-5007-2015, 2015.
- Engelmann, R., Kanitz, T., Baars, H., Heese, B., Althausen, D., Skupin, A., Wandinger, U., Komppula, M., Stachlewska, I. S., Amiridis, V., Marinou, E., Mattis, I., Linné, H., and Ansmann, A.: The automated multiwavelength Raman polarization and water-vapor lidar Polly<sup>XT</sup>: the neXT generation, *Atmos. Meas. Tech.*, 9, 1767–1784, doi:10.5194/amt-9-1767-2016, 2016.

- Fomba, K. W., Müller, K., van Pinxteren, D., Poulain, L., van Pinxteren, M., and Herrmann, H.: Long-term chemical characterization of tropical and marine aerosols at the Cape Verde Atmospheric Observatory (CVAO) from 2007 to 2011, *Atmos. Chem. Phys.*, 14, 8883–8904, doi:10.5194/acp-14-8883-2014, 2014.
- Fraser, R. S.: Optical thickness of atmospheric dust over Tadjikistan, *Atmos. Environ.*, 27, 2533–2538, doi:10.1016/0960-1686(93)90026-U, 1993.
- Freedman, E. and Neuzil, M., eds.: *Environmental Crises in Central Asia: From steppes to seas, from deserts to glaciers*, Routledge Studies in Environmental Communication and Media, Taylor & Francis, 2015.
- Freudenthaler, V.: About the effects of polarising optics on lidar signals and the  $\Delta 90$  calibration, *Atmos. Meas. Tech.*, 9, 4181–4255, doi:10.5194/amt-9-4181-2016, 2016.
- Freudenthaler, V., Esselborn, M., Wiegner, M., Heese, B., Tesche, M., Ansmann, A., Müller, D., Althausen, D., Wirth, M., Fix, A., Ehret, G., Knippertz, P., Toledano, C., Gasteiger, J., Garhammer, M., and Seefeldner, M.: Depolarization ratio profiling at several wavelengths in pure Saharan dust during SAMUM 2006, *Tellus B*, 61, 165–179, doi:10.1111/j.1600-0889.2008.00396.x, 2009.
- Gabbi, J., Huss, M., Bauder, A., Cao, F., and Schwikowski, M.: The impact of Saharan dust and black carbon on albedo and long-term mass balance of an Alpine glacier, *Cryosphere*, 9, 1385–1400, doi:10.5194/tc-9-1385-2015, 2015.
- GDAS: Global Data Assimilation System, meteorological data base, available at: <https://www.ready.noaa.gov/gdas1.php>, last access: February, 2017.
- Georgoulas, A. K., Alexandri, G., Kourtidis, K. A., Lelieveld, J., Zanis, P., Pöschl, U., Levy, R., Amiridis, V., Marinou, E., and Tsikerdekis, A.: Spatiotemporal variability and contribution of different aerosol types to the aerosol optical depth over the Eastern Mediterranean, *Atmos. Chem. Phys.*, 16, 13 853–13 884, doi:10.5194/acp-16-13853-2016, 2016.
- Giannakaki, E., van Zyl, P. G., Müller, D., Balis, D., and Komppula, M.: Optical and microphysical characterization of aerosol layers over South Africa by means of multi-wavelength depolarization and Raman lidar measurements, *Atmos. Chem. Phys.*, 16, 8109–8123, doi:10.5194/acp-16-8109-2016, 2016.
- Gillette, D. A. and Dobrowolski, J. P.: Soil crust formation by dust deposition at Shaartuz, Tadjik, S.S.R., *Atmos. Environ.*, 27, 2519–2525, doi:10.1016/0960-1686(93)90024-S, 1993.
- Gillette, D. A., Bodhaine, B. A., and D, M.: Transport and deposition of desert dust in the Kafirnigan River Valley (Tadjikistan) from Shaartuz to Esanbay: Measurements and a simple model, *Atmos. Environ.*, 27, 2545–2552, doi:10.1016/0960-1686(93)90028-W, 1993.
- Gimmetstad, G. G.: Reexamination of depolarization in lidar measurements, *Appl. Opt.*, 47, 3795–3802, doi:10.1364/AO.47.003795, 2008.
- Gkikas, A., Hatzianastassiou, N., and Mihalopoulos: Aerosol events in the broader Mediterranean basin based on 7-year (2000–2007) MODIS C005 data, *Ann. Geophys.*, 27, 3509–3522, doi:10.5194/angeo-27-3509-2009, 2009.
- Gkikas, A., Hatzianastassiou, N., Mihalopoulos, N., Katsoulis, V., Kazadzis, S., Pey, J., Querol, X., and Torres, O.: The regime of intense desert dust episodes in the Mediterranean based on contemporary satellite observations and ground measurements, *Atmos. Chem. and Phys.*, 13, 12 135–12 154, doi:10.5194/acp-13-12135-2013, 2013.
- Gkikas, A., Basart, S., Hatzianastassiou, N., Marinou, E., Amiridis, V., Kazadzis, S., Pey, J., Querol, X., Jorba, O., Gassó, S., and Baldasano, J. M.: Mediterranean intense desert dust outbreaks and their vertical structure based on remote sensing data, *Atmos. Chem. Phys.*, 16, 8609–8642, doi:10.5194/acp-16-8609-2016, 2016.
- Golitsyn, G. and Gillette, D. A.: Introduction: A joint Soviet-American experiment for the study of Asian desert dust and its impact on local meteorological conditions and climate, *Atmos. Environ.*, 27, 2467–2470, doi:10.1016/0960-1686(93)90017-S, 1993.

- Golitsyn, G. S., Gillette, D. A., and Johnson, T., eds.: Joint Soviet-American Experiment on Arid Aerosol, Hydrometeoizdat, St. Petersburg, 1993a.
- Golitsyn, G. S., Shukurov, A. K., F., A. S., and I, N. B.: On the surface air cooling due too dust atmospheric turbidity, pp. 67–78, in: Golitsyn et al. (1993a), 1993b.
- 5 Gomes, L. and Gillette, D. A.: A comparison of characteristics of aerosol from dust storms in Central Asia with soil-derived dust from other regions, *Atmos. Environ.*, 27, 2539–2544, doi:10.1016/0960-1686(93)90027-V, 1993.
- Griffin, D. W.: Atmospheric Movement of Microorganisms in Clouds of Desert Dust and Implications for Human Health, *Clin. Microbiol. Rev.*, 20, 459–477, doi:10.1128/CMR.00039-06, 2007.
- Groll, M., Opp, C., and Aslanov, I.: Spatial and temporal distribution of the dust deposition in Central Asia – results from a long term  
10 monitoring program, *Aeolian Res.*, 9, 49–62, doi:10.1016/j.aeolia.2012.08.002, 2013.
- Gross, B., Moshary, F., and Arend, M., eds.: EPJ Web of Conferences Volume 119 (2016): The 27th International Laser Radar Conference (ILRC 27), 2016.
- Groß, S., Tesche, M., Freudenthaler, V., Toledano, C., Wiegner, M., Ansmann, A., Althausen, D., and Seefeldner, M.: Characterization of  
15 Saharan dust, marine aerosols and mixtures of biomass-burning aerosols and dust by means of multi-wavelength depolarization and Raman lidar measurements during SAMUM 2, *Tellus B*, 64, 706–724, doi:10.1111/j.1600-0889.2011.00556.x, 2011.
- Groß, S., Freudenthaler, V., Schepanski, K., Toledano, C., Schäfler, A., Ansmann, A., and Weinzierl, B.: Optical properties of long-range transported Saharan dust over Barbados as measured by dual-wavelength depolarization Raman lidar measurements, *Atmos. Chem. Phys.*, 15, 11 067–11 080, doi:10.5194/acp-15-11067-2015, 2015.
- Haarig, M., Althausen, D., Ansmann, A., Klepel, A., Baars, H., Engelmann, R., Groß, S., and Freudenthaler, V.: Measurement of the Linear  
20 Depolarization Ratio of Aged Dust at Three Wavelengths (355, 532 and 1064 nm) Simultaneously over Barbados, in: Gross et al. (2016), p. 18009, doi:10.1051/epjconf/201611918009, 2016.
- Haarig, M., Ansmann, A., Althausen, D., Klepel, A., Groß, S., Freudenthaler, V., Toledano, C., Mamouri, R.-E., Farrell, D. A., Prescod, D. A., Marinou, E., Burton, S. P., Gasteiger, J., Engelmann, R., and Baars, H.: Triple-wavelength depolarization-ratio profiling of Saharan dust over Barbados during SALTRACE in 2013 and 2014, *Atmos. Chem. Phys. Discuss.*, 2017, 1–43, doi:10.5194/acp-2017-170, 2017.
- 25 Hansen, A. D. A., Kapustin, V. A., Kopeikin, V. M., Gillette, D. A., and Bodhaine, B. A.: Optical absorption by aerosol black carbon and dust in a desert region of Central Asia, *Atmos. Environ.*, 27, 2527–2531, doi:10.1016/0960-1686(93)90025-T, 1993.
- Hara, K. and Zhang, D.: Bacterial abundance and viability in long-range transported dust, *Atmos. Env.*, 47, 20–25, doi:10.1016/j.atmosenv.2011.11.050, 2012.
- Heinold, B., Tegen, I., Bauer, S., and Wendisch, M.: Regional modelling of Saharan dust and biomass-burning smoke Part 2: Direct radiative  
30 forcing and atmospheric dynamic response, *Tellus B*, 63, doi:10.1111/j.1600-0889.2011.00574.x, 2011.
- Heinold, B., Tegen, I., Schepanski, K., and Banks, J. R.: New developments in the representation of Saharan dust sources in the aerosol–climate model ECHAM6-HAM2, *Geosci. Model Dev.*, 9, 765–777, doi:10.5194/gmd-9-765-2016, 2016.
- Heintzenberg, J.: The SAMUM-1 experiment over Southern Morocco: overview and introduction, *Tellus B*, 61, 2–11, doi:10.1111/j.1600-0889.2008.00403.x, 2009.
- 35 Holben, B. N., Eck, T. F., Slutsker, I., Tanré, D., Buis, J. P., Setzer, A., E., V., Reagan, J. A., Kaufman, Y. J., Nakajima, T., Lavenu, F., Jankowiak, I., and Smirnov, A.: AERONET—A Federated Instrument Network and Data Archive for Aerosol Characterization, *Rem. Sens. Environ.*, 66, 1–16, doi:10.1016/S0034-4257(98)00031-5, 1998.

- Iwasaka, Y., Shi, G.-Y., Yamada, M., Matsuki, A., Trochkin, D., Kim, Y. S., Zhang, D., Nagatani, T., Shibata, T., Nagatani, M., Nakata, H., Shen, Z., Li, G., and Chen, B.: Importance of dust particles in the free troposphere over the Taklamakan Desert: Electron microscopic experiments of particles collected with a balloonborne particle impactor at Dunhuang, China, *J. Geophys. Res. Atmos.*, 108, 8644, doi:10.1029/2002JD003270, 2003.
- 5 Jin, Y., Kai, K., Shibata, T., Zhang, K., and Zhou, H.: Validation of the Dust Layer Structure over the Taklimakan Desert, China by the CALIOP Space-Borne Lidar Using Ground-Based Lidar, *SOLA*, 6, 121–124, doi:10.2151/sola.2010-031, 2010.
- Kai, K., Nagata, Y., Tsunematsu, N., Matsumura, T., Kim, H.-S., Matsumoto, T., Hu, S., Zhou H., Abo, M., and Nagai, T.: The structure of the dust layer over the Taklimakan Desert during the dust storm in April 2002 as observed using a depolarization lidar, *J. Meteor. Soc. Japan*, 86, 1–16, doi:10.2151/jmsj.86.1, 2008.
- 10 Kandler, K., Schütz, L., Deutscher, C., Ebert, M., Hofmann, H., Jäckel, S., Jaenicke, R., Knippertz, P., Lieke, K., Massling, A., Petzold, A., Schladitz, A., Weinzierl, B., Wiedensohler, A., Zorn, S., and Weinbruch, S.: Size distribution, mass concentration, chemical and mineralogical composition and derived optical parameters of the boundary layer aerosol at Tinfou, Morocco, during SAMUM 2006, *Tellus B*, 61, 32–50, doi:10.1111/j.1600-0889.2008.00385.x, 2009.
- Kandler, K., Lieke, K., Benker, N., Emmel, C., Küpper, M., Müller-Ebert, D., Ebert, M., Scheuven, D., Schladitz, A., Schütz, L., and
- 15 Weinbruch, S.: Electron microscopy of particles collected at Praia, Cape Verde, during the Saharan Mineral Dust Experiment: particle chemistry, shape, mixing state and complex refractive index, *Tellus B*, 63, 475–496, doi:10.1111/j.1600-0889.2011.00550.x, 2011.
- Kanitz, T., Ansmann, A., Engelmann, R., and Althausen, D.: North-south cross sections of the vertical aerosol distribution over the Atlantic Ocean from multiwavelength Raman/polarization lidar during Polarstern cruises, *J. Geophys. Res. Atmos.*, 118, 2643–2655, doi:10.1002/jgrd.50273, 2013.
- 20 Kayumov, A.: Glaciers Resources of Tajikistan in Condition of the Climate Change, State Agency for Hydrometeorology of Committee for Environmental Protection under the Government of the Republic of Tajikistan, 2010.
- Kayumov, A. K. and Novikov, V. V.: The Third National Communication of the Republic of Tajikistan under the United Nations Framework Convention on Climate Change, Tech. rep., The State Agency for Hydrometeorology, Committee for environmental protection under the of the Republic of Tajikistan, Dushanbe, [http://unfccc.int/resource/docs/natc/tjknc3\\_eng.pdf](http://unfccc.int/resource/docs/natc/tjknc3_eng.pdf), 2014.
- 25 Klüser, L., Martynenko, D., Banks, J. R., Bergemann, C., Brindley, H. E., and Holzer-Popp, T.: Information content of space-borne hyperspectral infrared observations with respect to mineral dust properties, *Remote Sens. Environ.*, 156, 294–309, doi:10.1016/j.rse.2014.09.036, 2015.
- Lioubimtseva, E. and Henebry, G. M.: Climate and environmental change in arid Central Asia: Impacts, vulnerability, and adaptations, *J. Arid Environ.*, 73, 963–977, doi:10.1016/j.jaridenv.2009.04.022, 2009.
- 30 Liu, Z., Sugimoto, N., and Murayama, T.: Extinction-to-backscatter ratio of Asian dust observed with high-spectral-resolution lidar and Raman lidar, *Appl. Opt.*, 41, 2760–2767, doi:10.1364/AO.41.002760, 2002.
- Makhmaliev, B. and Novikov, V. V.: The First National Communication of the Republic of Tajikistan under the United Nations Framework Convention on Climate Change, Tech. rep., Ministry for Nature Protection of the Republic of Tajikistan, The Main Administration on Hydrometeorology and Environmental Monitoring, Dushanbe, <http://unfccc.int/resource/docs/natc/tajnc1.pdf>, 2002.
- 35 Makhmaliev, B. U., Novikov, V. V., and Kayumov, A. K.: The First National Communication of the Republic of Tajikistan under the United Nations Framework Convention on Climate Change, Phase 2, Tech. rep., Ministry for Nature Protection of the Republic of Tajikistan, The Main Administration on Hydrometeorology and Environmental Monitoring, Dushanbe, <http://unfccc.int/resource/docs/natc/tajnc1add.pdf>, 2004.

- Makhmadaliev, B. U., Kayumov, A. K., Novikov, V. V., Mustaeva, N., and Rajobov, I.: The Second National Communication of the Republic of Tajikistan under the United Nations Framework Convention on Climate Change, Tech. rep., The State Agency for Hydrometeorology, Committee for environmental protection under the of the Republic of Tajikistan, Dushanbe, <http://unfccc.int/resource/docs/natc/tainc2.pdf>, 2008.
- 5 Mamouri, R. E. and Ansmann, A.: Fine and coarse dust separation with polarization lidar, *Atmos. Meas. Tech.*, 7, 3717–3735, doi:10.5194/amt-7-3717-2014, 2014.
- Mamouri, R. E. and Ansmann, A.: Potential of polarization/Raman lidar to separate fine dust, coarse dust, maritime, and anthropogenic aerosol profiles, *Atmos. Meas. Tech. Discuss.*, 2017, 1–41, doi:10.5194/amt-2017-131, 2017.
- Mamouri, R. E., Ansmann, A., Nisantzi, A., Kokkalis, P., Schwarz, A., and Hadjimitsis, D. G.: Low Arabian dust extinction-to-backscatter ratio, *Geophys. Res. Lett.*, 40, 4762–4766, doi:10.1002/grl.50898, 2013.
- 10 Mamouri, R. E., Ansmann, A., Nisantzi, A., Solomos, S., Kallos, G., and Hadjimitsis, D. G.: Extreme dust storm over the eastern Mediterranean in September 2015: satellite, lidar, and surface observations in the Cyprus region, *Atmos. Chem. Phys.*, 16, 13 711–13 724, doi:10.5194/acp-16-13711-2016, 2016.
- Mikami, M., Shi, G. Y., Uno, I., Yabuki, S., Iwasaka, Y., Yasui, M., Aoki, T., Tanaka, T. Y., Kurosaki, Y., Masuda, K., Uchiyama, A., Matsuki, A., Sakai, T., Takemi, T., Nakawo, M., Seino, N., Ishizuka, M., Satake, S., Fujita, K., Hara, Y., Kai, K., Kanayama, S., Hayashi, M., Du, M., Kanai, Y., Yamada, Y., Zhang, X. Y., Shen, Z., Zhou, H., Abe, O., Nagai, T., Tsutsumi, Y., Chiba, M., and Suzuk, J.: Aeolian dust experiment on climate impact: An overview of Japan China joint project ADEC, *Glob. Planet. Change*, 52, 142–172, doi:10.1016/j.gloplacha.2006.03.001, 2006.
- 15 Miller-Schulze, J. P., , Shafer, M. M., Schauer, J. J., A., S. P., Lantz, J., Artamonova, M., Chen, B. Imashev, S., Sverdlik, L., Carmichael, G. R., and Deminter, J. T.: Characteristics of fine particle carbonaceous aerosol at two remote sites in Central Asia, *Atmos. Environ.*, 45, 695–6964, doi:10.1016/j.atmosenv.2011.09.026, 2011.
- 20 Müller, D., Ansmann, A., Mattis, I., Tesche, M., Wandinger, U., Althausen, D., and Pisani, G.: Aerosol-type-dependent lidar ratios observed with Raman lidar, *J. Geophys. Res. Atmos.*, 112, doi:10.1029/2006JD008292, 2007.
- Nazarov, B. I., Abdullaev, S. F., , and Maslov, V. A.: Studies of Temperature Effects of Dust Storms, *Izv. Atmos. Ocean. Phys.*, 46, 475–481, 2010a.
- 25 Nazarov, B. I., Maslov, V. A., , and F., A. S.: Optical and Microphysical Parameters of Arid Dust Aerosol, *Izv. Atmos. Ocean. Phys.*, 46, 468–474, 2010b.
- Nisantzi, A., Mamouri, R. E., Ansmann, A., Schuster, G. L., and Hadjimitsis, D. G.: Middle East versus Saharan dust extinction-to-backscatter ratios, *Atmos. Chem. Phys.*, 15, 5203–5240, doi:10.5194/acpd-15-5203-2015, 2015.
- 30 Nobakht, M., Shahgedanova, M., and White, K.: Characterization of dust sources in Central Asia using Remote Sensing, AGU Fall Meeting Abstracts, 2015.
- Ott, D. K. and Peters, T.: A Shelter to Protect a Passive Sampler for Coarse Particulate Matter, PM10-2.5, *Aerosol Sci. Technol.*, 42, 299–309, doi:10.1080/02786820802054236, 2008.
- Pachenko, M. V., Terpugova, S. A., Bodhaine, B. A., Isakov, A. A., Sviridenkov, M. A., Sokolik, I. N., Romashova, E. V., Nazarov, B. I., Shukurov, A. K., Chistyakova, E. I., and Johnson, T. C.: Optical investigations of dust storms during U.S.S.R.-U.S. experiments in Tadzhikistan, 1989, *Atmos. Environ.*, 27, 2503–2508, doi:10.1016/0960-1686(93)90023-R, 1993.
- 35 Park, J., Ichijo, T., Nasu, M., and Yamaguchi, N.: Investigation of bacterial effects of Asian dust events through comparison with seasonal variability in outdoor airborne bacterial community, *Sci. Rep.*, 6, 35706, doi:10.1038/srep35706, 2016.



- Ridley, D. A., Heald, C. L., Kok, J. F., and Zhao, C.: An observationally constrained estimate of global dust aerosol optical depth, *Atmos. Chem. Phys.*, 16, 15 097–15 117, doi:10.5194/acp-16-15097-2016, 2016.
- Rolph, G.: Real-time Environmental Applications and Display sYstem (READY) Website (<http://ready.arl.noaa.gov>), NOAA Air Resources Laboratory, Silver Spring, MD, 2016.
- 5 Ryder, C. L., McQuaid, J. B., Flamant, C., Rosenberg, P. D., Washington, R., Brindley, H. E., Highwood, E. J., Marsham, J. H., Parker, D. J., Todd, M. C., Banks, J. R., Brooke, J. K., Engelstaedter, S., Estelles, V., Formenti, P., Garcia-Carreras, L., Kocha, C., Marengo, F., Sodemann, H., Allen, C. J. T., Bourdon, A., Bart, M., Cavazos-Guerra, C., Chevaillier, S., Crosier, J., Darbyshire, E., Dean, A. R., Dorsey, J. R., Kent, J., O'Sullivan, D., Schepanski, K., Szpek, K., Trembath, J., and Woolley, A.: Advances in understanding mineral dust and boundary layer processes over the Sahara from Fennec aircraft observations, *Atmos. Chem. Phys.*, 15, 8479–8520, doi:10.5194/acp-15-8479-2015, 2015.
- 10 Sakai, T., Nagai, T., Nakazato, M., Mano, Y., and Matsumura, T.: Ice clouds and Asian dust studied with lidar measurements of particle extinction-to-backscatter ratio, particle depolarization, and water-vapor mixing ratio over Tsukuba, *Appl. Opt.*, 42, 7103–7116, doi:10.1364/AO.42.007103, 2003.
- Schettler, G., Shabunin, A., Kemnitz, H., Knoeller, K., Imashev, S., Rybin, A., and Wetzol, H.-U.: Seasonal and diurnal variations in dust characteristics on the northern slopes of the Tien Shan – Grain-size, mineralogy, chemical signatures and isotope composition of attached nitrate, *J. Asian Earth Sci.*, 88, 257–276, doi:10.1016/j.jseae.2014.03.019, 2014.
- 15 Schmale, J., Flanner, M., Kang, S., Sprenger, M., Zhang, Q., Guo, J., Yang Li, Y., Schwikowski, M., and Farinotti, D.: Modulation of snow reflectance and snowmelt from Central Asian glaciers by anthropogenic black carbon, *Sci. Rep.*, 7, doi:10.1038/srep40501, 2017.
- Schuster, G. L., Vaughan, M., MacDonnell, D., Su, W., Winker, D., Dubovik, O., Lapyonok, T., and Trepte, C.: Comparison of CALIPSO aerosol optical depth retrievals to AERONET measurements, and a climatology for the lidar ratio of dust, *Atmos. Chem. Phys.*, 12, 7431–7452, doi:10.5194/acp-12-7431-2012, 2012.
- 20 Schwarz, A.: Aerosol typing over Europe and its benefits for the CALIPSO and EarthCARE missions - Statistical analysis based on multiwavelength aerosol lidar measurements from ground-based EARLINET stations and comparison to spaceborne CALIPSO data, Ph.D. thesis, University of Leipzig, 2015.
- 25 Semenov, V. K., Smirnov, A., Aref'ev, V. N., Sinyakov, V. P., Sorokina, L. I., and Ignatova, N. I.: Aerosol optical depth over the mountainous region in central Asia (Issyk-Kul Lake, Kyrgyzstan), *Geophys. Res. Lett.*, 32, doi:10.1029/2004GL021746, 105807, 2005.
- Shukurov, A. K., Nazarov, B. I., Abdullaev, S. F., and V., P. S.: On optical depth ratios of dust aerosol in visible and infrared spectra regions, pp. 83–88, in: Golitsyn et al. (1993a), 1993.
- Siegfried, T., Bernauer, T., Guennet, R., Sellars, S., Robertson, A. W., Mankin, J., Bauer-Gottwein, P., and Yakovlev, A.: Will climate change exacerbate water stress in Central Asia?, *Clim. Change*, 112, 881–899, doi:10.1007/s10584-011-0253-z, 2012.
- 30 Smirnov, V. V., Johnson, T. C., Krapivtseva, G. M., Krivchikova, T. V., and Shukurov, A. H.: Synoptic meteorological conditions during the U.S.S.R./U.S. dust experiment in Tadzhikistan in September 1989, *Atmos. Environ.*, 27, 2471–2479, doi:10.1016/0960-1686(93)90018-T, 1993.
- Smirnov, V. V., Gillette, D. A., Golitsyn, G. S., and MacKinnon, D. J.: The origin and evolution of dust clouds in Central Asia, *Atmos. Res.*, 34, 169–176, doi:10.1016/0169-8095(94)90088-4, 1994.
- 35 Sokolik, I. N. and Golitsyn, G.: Investigation of optical and radiative properties of atmospheric dust aerosols, *Atmos. Environ.*, 27, 2509–2517, doi:10.1016/0960-1686(93)90021-P, 1993.

- Sokolik, I. N., Andronova, A. V., and Johnson, T. C.: Complex refractive index of atmospheric dust aerosols, *Atmos. Environ.*, 27, 2495–2502, doi:10.1016/0960-1686(93)90021-P, 1993.
- Sorg, A., B., M., G., S., A., A., M., H. C., and M., S.: Coping with changing water resources: The case of the Syr Darya river basin in Central Asia, *Environ. Sci. Policy*, 43, 68–77, doi:https://doi.org/10.1016/j.envsci.2013.11.003, 2014.
- 5 Stein, A. F., Draxler, R. R., Rolph, G. D., Stunder, B. J. B., Cohen, M. D., and Ngan, F.: NOAA’s HYSPLIT atmospheric transport and dispersion modeling system, *Bull. Amer. Meteor. Soc.*, 96, 2059–2077, doi:10.1175/BAMS-D-14-00110.1, 2015.
- Stohl, A., Forster, C., Frank, A., Seibert, P., and Wotawa, G.: Technical note: The Lagrangian particle dispersion model FLEXPART version 6.2., *Atmos. Chem. Phys.*, 5, 2461–2474, doi:10.5194/acp-5-2461-2005, 2005.
- Su, L. and Toon, O. B.: Saharan and Asian dust: similarities and differences determined by CALIPSO, AERONET, and a coupled climate-aerosol microphysical model, *Atmos. Chem. Phys.*, 11, 3263–3280, doi:10.5194/acp-11-3263-2011, 2011.
- 10 Tanaka, T. Y., Kurosaki, Y., Chiba, M., Matsumura, T., Nagai, T. and Yamazaki, A., Uchiyama, A., Tsunematsu, N., and Kai, K.: Possible transcontinental dust transport from North Africa and the Middle East to East Asia, *Atmos. Environ.*, 39, 3901–3909, doi:10.1016/j.atmosenv.2005.03.034, 2005.
- Tesche, M., Ansmann, A., Müller, D., Althausen, D., Engelmann, R., Hu, M., and Zhang, Y.: Particle backscatter, extinction, and lidar ratio profiling with Raman lidar in south and north China, *Appl. Opt.*, 46, 6302–6308, doi:10.1364/AO.46.006302, 2007.
- 15 Tesche, M., Ansmann, A., Müller, D., Althausen, D., Engelmann, R., Freudenthaler, V., and Groß, S.: Vertically resolved separation of dust and smoke over Cape Verde by using multiwavelength Raman and polarization lidars during Saharan Mineral Dust Experiment 2008, *J. Geophys. Res. Atmos.*, 114, doi:10.1029/2009JD011862, 2009a.
- Tesche, M., Ansmann, A., Müller, D., Althausen, D., Mattis, I., Heese, B., Freudenthaler, V., Wiegner, M., Esselborn, M., Pisani, G., and Knippertz, P.: Vertical profiling of Saharan dust with Raman lidars and airborne HSRL in southern Morocco during SAMUM, *Tellus B*, 20 61, 144–164, doi:10.1111/j.1600-0889.2008.00390.x, 2009b.
- Tesche, M., Wandinger, U., Ansmann, A., Althausen, D., Müller, D., and Omar, A. H.: Ground-based validation of CALIPSO observations of dust and smoke in the Cape Verde region, *J. Geophys. Res. Atmos.*, 118, 2889–2902, doi:10.1002/jgrd.50248, 2013.
- Tsunematsu, N., Sato, T., Kimura, F., Kai, K., Kurosaki, Y., Nagai, T., Zhou, H., and Mikami, M.: Extensive dust outbreaks following the morning inversion breakup in the Taklimakan Desert, *J. Geophys. Res. Atmos.*, 110, doi:10.1029/2005JD005994, d21207, 2005.
- 25 Uno, I., Eguchi, K., Yumimoto, K., Takemura, T., Shimizu, A., Uematsu, M., Liu, Z., Wang, Z., Hara, Y., and Sugimoto, N.: Asian dust transported one full circuit around the globe, *Nat. Geosci.*, 2, 557–560, doi:10.1038/ngeo583, 2009.
- Veselovskii, I., Goloub, P., Podvin, T., Bovchaliuk, V., Derimian, Y., Augustin, P., Fourmentin, M., Tanre, D., Korenskiy, M., Whiteman, D. N., Diallo, A., Ndiaye, T., Kolgotin, A., and Dubovik, O.: Retrieval of optical and physical properties of African dust from multiwavelength Raman lidar measurements during the SHADOW campaign in Senegal, *Atmos. Chem. Phys.*, 16, 7013–7028, doi:10.5194/acp-16-7013-2016, 2016.
- 30 Wandinger, U., Tesche, M., Seifert, P., Ansmann, A., Müller, D., and Althausen, D.: Size matters: Influence of multiple scattering on CALIPSO light-extinction profiling in desert dust, *Geophys. Res. Lett.*, 37, L10 801, doi:10.1029/2010GL042815, 2010.
- Weinzierl, B., Ansmann, A., Prospero, J. M., Althausen, D., Benker, N., Chouza, F., Dollner, M., Farrell, D., Fomba, W. K., Freudenthaler, V., Gasteiger, J., Groß, S., Haarig, M., Heinold, B., Kandler, K., Kristensen, T. B., Mayol-Bracero, O.-L., Müller, T., Reitebuch, O., Sauer, D., Schäfler, A., Schepanski, K., Spanu, A., Tegen, I., Toledano, C., and Walser, A.: The Saharan Aerosol Long-range TRansport and Aerosol Cloud Interaction Experiment (SALTRACE): overview and selected highlights, *Bull. Amer. Meteorol. Soc.*, 1, 1–2, doi:10.1175/BAMS-D-15-00142.1, 2017.

- Wiggs, G. F. S., O'hara, S. L., Wegerdt, J., Van Der Meer, J., Small, I., and Hubbard, R.: The dynamics and characteristics of aeolian dust in dryland Central Asia: possible impacts on human exposure and respiratory health in the Aral Sea basin, *Geogr. J.*, 169, 142–157, doi:10.1111/1475-4959.04976, 2003.
- Wolke, R., Schröder, W., Roland Schrödner, R., and Renner, E.: Influence of grid resolution and meteorological forcing on simulated European air quality: A sensitivity study with the modeling system COSMO–MUSCAT, *Atmos. Environ.*, 53, 110–30, doi:10.1016/j.atmosenv.2012.02.085, 2012.
- Xi, X. and Sokolik, I. N.: Dust interannual variability and trend in Central Asia from 2000 to 2014 and their climatic linkages, *J. Geophys. Res. Atmos.*, 120, doi:10.1002/2015JD024092, 2015a.
- Xi, X. and Sokolik, I. N.: Seasonal dynamics of threshold friction velocity and dust emission in Central Asia, *J. Geophys. Res. Atmos.*, 120, 1536–1564, doi:10.1002/2014JD02247, 2015b.
- Xi, X. and Sokolik, I. N.: Quantifying the anthropogenic dust emission from agricultural land use and desiccation of the Aral Sea in Central Asia, *J. Geophys. Res. Atmos.*, 121, doi:10.1002/2016JD025556, 2016.
- Xu, H., Wang, X., and Zhang, X.: Decreased vegetation growth in response to summer drought in Central Asia from 2000 to 2012, *Int. J. Appl. Earth Obs. Geoinf.*, 52, 390–402, doi:10.1016/j.jag.2016.07.010, 2016.
- Yamaguchi, N., Ichijo, T., Sakotani, A., Baba, T., and Nasu, M.: Global dispersion of bacterial cells on Asian dust, *Sci. Rep.*, 2, 525, doi:10.1038/srep00525, 2015.
- Zhang, K., O'Donnell, D., Kazil, J., Stier, P., Kinne, S., Lohmann, U., Ferrachat, S., Croft, B., Quaas, J., Wan, H., Rast, S., and Feichter, J.: The global aerosol-climate model ECHAM-HAM, version 2: sensitivity to improvements in process representations, *Atmos. Chem. and Phys.*, 12, 8911–8949, doi:10.5194/acp-12-8911-2012, 2012.

**Table 1.** Products of the Polly<sup>XT</sup> lidar and their relative uncertainties.  $\beta$ : particle backscatter coefficient,  $\alpha$ : particle extinction coefficient,  $S$ : lidar ratio,  $\tilde{\alpha}$ : extinction-related Ångström exponent,  $\delta$ : particle linear depolarization ratio, WV: water vapor, and dust and non-dust mass concentrations.

Product	Wavelengths [nm]	Uncertainty [%]
$\beta$	355, 532, 1064	5-10
$\alpha$	355 (from 387), 532 (from 607)	10-20
$S$	355, 532	11-22
$\tilde{\alpha}$	from $\alpha$ 355/532 pair	14-28
$\delta$	355, 532	7-11
WV	from 407	20
Mass conc.	from $\delta$ , $\beta$ , 355, 532	30-50

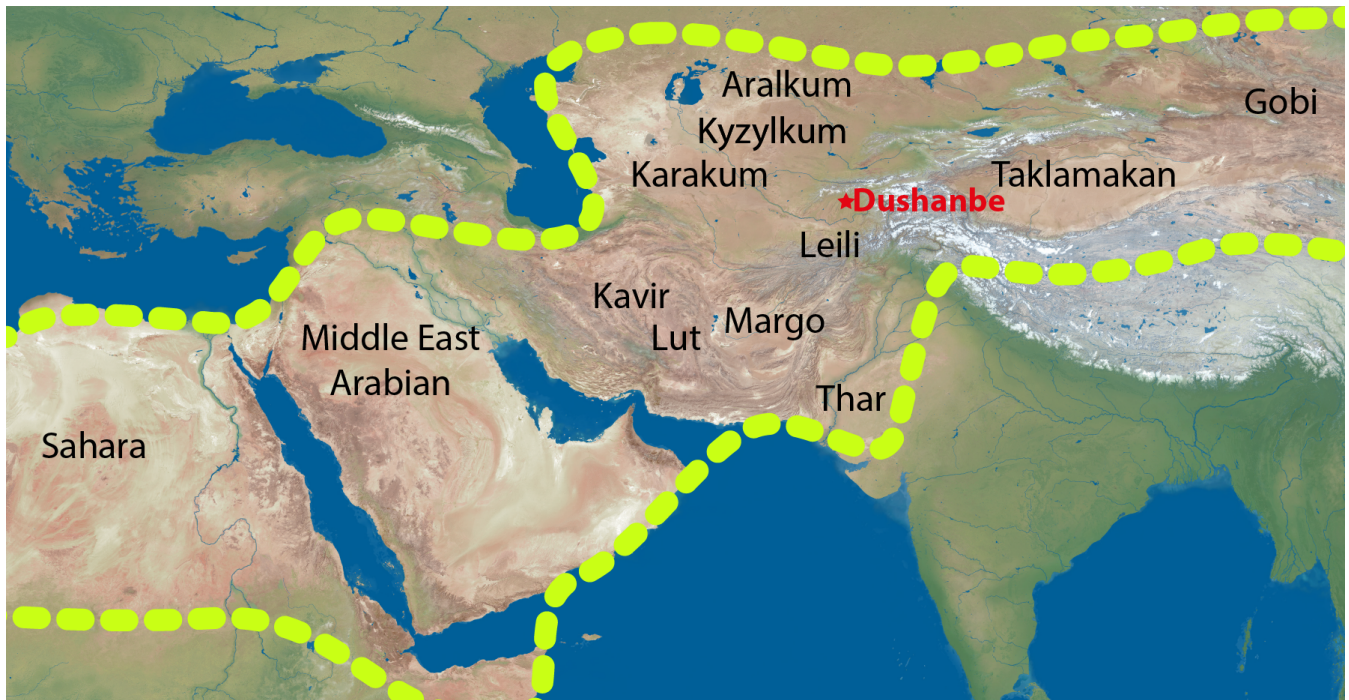
**Table 2.** Seasonal means of the sun photometer measurements (AERONET Level 1.5, cloudscreened). AOT at 500 nm wavelength, Ångström exponent (AE) from the 440-870 nm spectral range, and the fine mode fraction (FMF) from 500 nm wavelength.

Season	AOT	AE	FMF
Mar–May 2015	0.16	0.63	0.39
Jun–Aug 2015	0.40	0.54	0.33
Sep–Nov 2015	0.23	1.03	0.51
Dec–Feb 2015/16	0.15	1.28	0.71
Mar–May 2016	0.22	0.82	0.48
Jun–Aug 2016	0.41	0.52	0.34

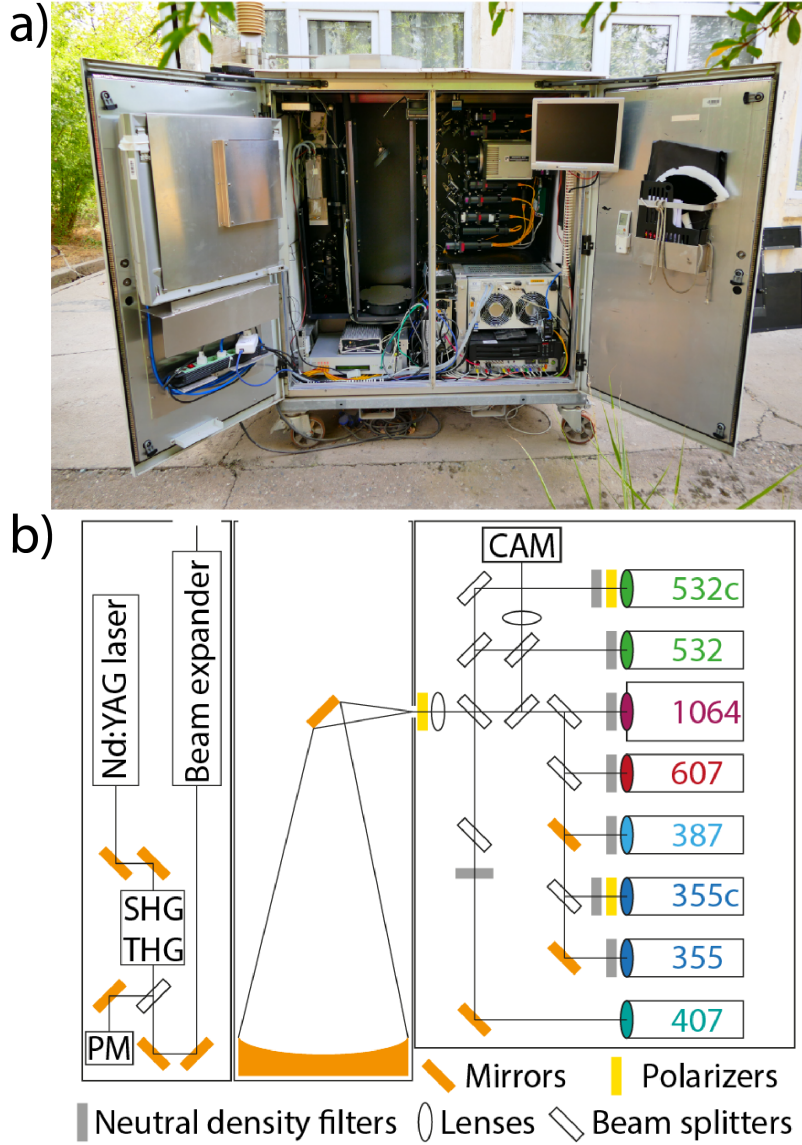
**Table 3.** Overview over the selected example cases.  $S_{355}, S_{532}$ : Lidar ratio at 355 nm and 532 nm wavelength.  $\delta_{355}^p, \delta_{532}^p$ : Particle linear depolarization ratio at 355 nm and 532 nm wavelength.  $\tau_{lidar}$ : Aerosol optical thickness measured with lidar at 532 nm wavelength.  $\tau_{SPM}$ : Aerosol optical thickness measurement of the co-located sun photometer at 500 nm wavelength (time of measurements in the footnote).  $\hat{\alpha}_{355/532}$ : Extinction-related Ångström exponent based on the particle extinction coefficients at 355 nm and 532 nm wavelength.  $\hat{\beta}_{355/532}, \hat{\beta}_{532/1064}$ : Backscatter-related Ångström exponents based on the particle backscatter coefficients at 355 nm, 532 nm, and 1064 nm wavelength. The height range to average the optical properties of the dust layers is placed within the core of each layer. The provided uncertainties are the standard deviations of the averaged values.

Case	Case 1	Case 2	Case 3	Case 4
	Lofted dust	Extreme dust	Most extreme dust	Pollution
Source region	Middle East	Central Asia	Central Asia	Local
Date	13 April 2015	8 August 2015	14 July 2016	13 May 2015
Time [UTC]	15:10–16:09	22:20–23:57	16:00–21:59	18:10–22:00
Layer height [km AGL]	2.3–4.9	0–3.1	0–2.6	0–4.1
$S_{355}$ [sr]	$42.2 \pm 3.0$	$46.9 \pm 2.1$	$40.3 \pm 0.6$	$55.8 \pm 7.1$
$S_{532}$ [sr]	$35.7 \pm 1.7$	$42.9 \pm 3.2$	$38.7 \pm 1.0$	$32.8 \pm 6.4$
$\delta_{355}^p$	$0.18 \pm 0.02$	$0.23 \pm 0.01$	$0.29 \pm 0.01$	$0.03 \pm 0.01$
$\delta_{532}^p$	$0.31 \pm 0.01$	$0.35 \pm 0.01$	$0.35 \pm 0.01$	$0.08 \pm 0.01$
$\tau_{lidar}$	0.40 (0–5.4 km)	1.50 (0–4.1 km)	3.89 (0–3 km)	0.13 (0–4.5 km)
$\tau_{SPM}$	0.41 <sup>(1)</sup>	1.19 <sup>(2)</sup> , 1.69 <sup>(3)</sup>	4.45 <sup>(4)</sup> , 2.89 <sup>(5)</sup>	0.18 <sup>(6)</sup>
$\hat{\alpha}_{355/532}$	$0.41 \pm 0.24$	$0.12 \pm 0.16$	$-0.08 \pm 0.06$	$2.07 \pm 0.72$
$\hat{\beta}_{355/532}$	$0.00 \pm 0.22$	$-0.11 \pm 0.09$	$-0.20 \pm 0.13$	$0.72 \pm 0.16$
$\hat{\beta}_{532/1064}$	$0.12 \pm 0.04$	$0.32 \pm 0.07$	$0.29 \pm 0.03$	$0.71 \pm 0.05$

Times of sun photometer measurements [UTC]: <sup>(1)</sup> 13 April 2015, 12:54:46, <sup>(2)</sup> 8 August 2015, 13:04:24, <sup>(3)</sup> 9 August 2015, 02:08:05, <sup>(4)</sup> 14 July 2016, 09:16:02, <sup>(5)</sup> 15 July 2016, 03:29:34, <sup>(6)</sup> 13 May 2015, 13:21:50

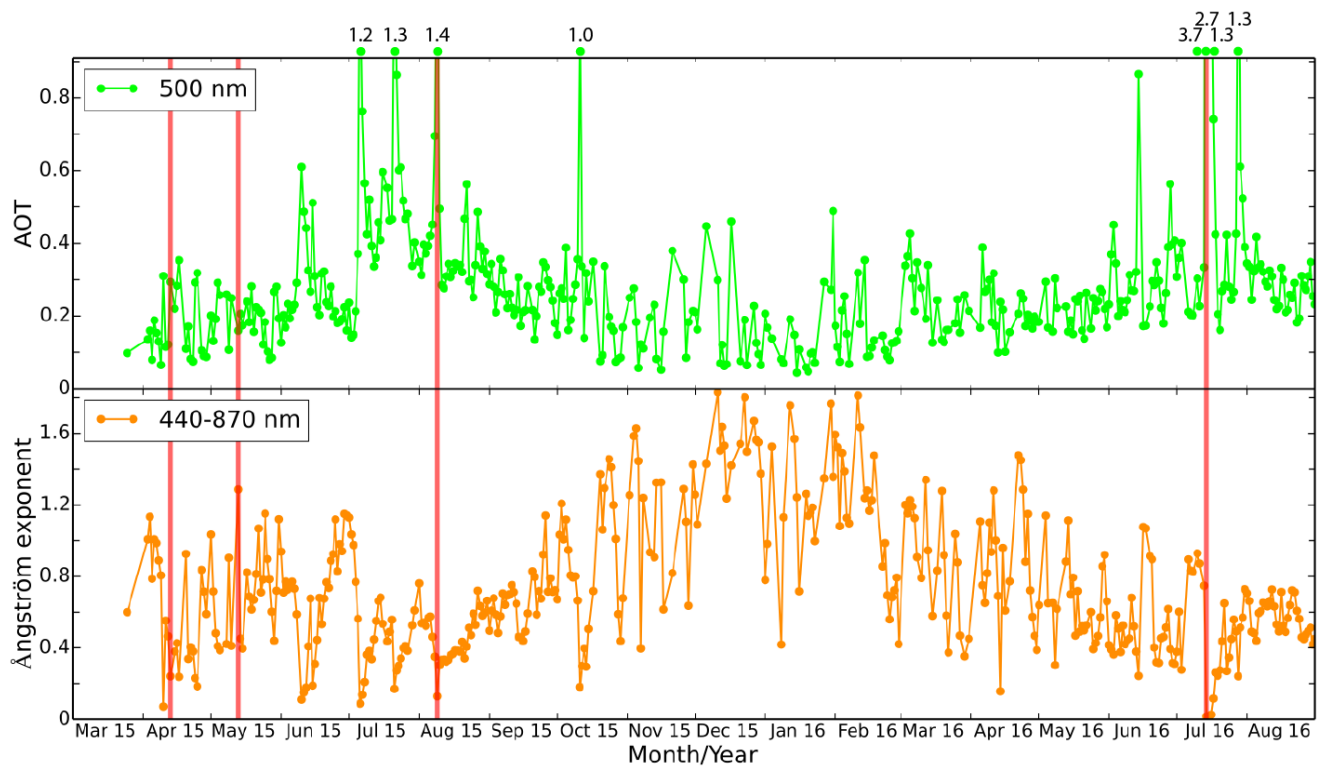


**Figure 1.** Geographical map with desert regions (black text), dust belt (bright green dotted line), measurement site in Dushanbe, Tajikistan (red text) highlighted (<http://naturalearth.springercarto.com>, adapted).

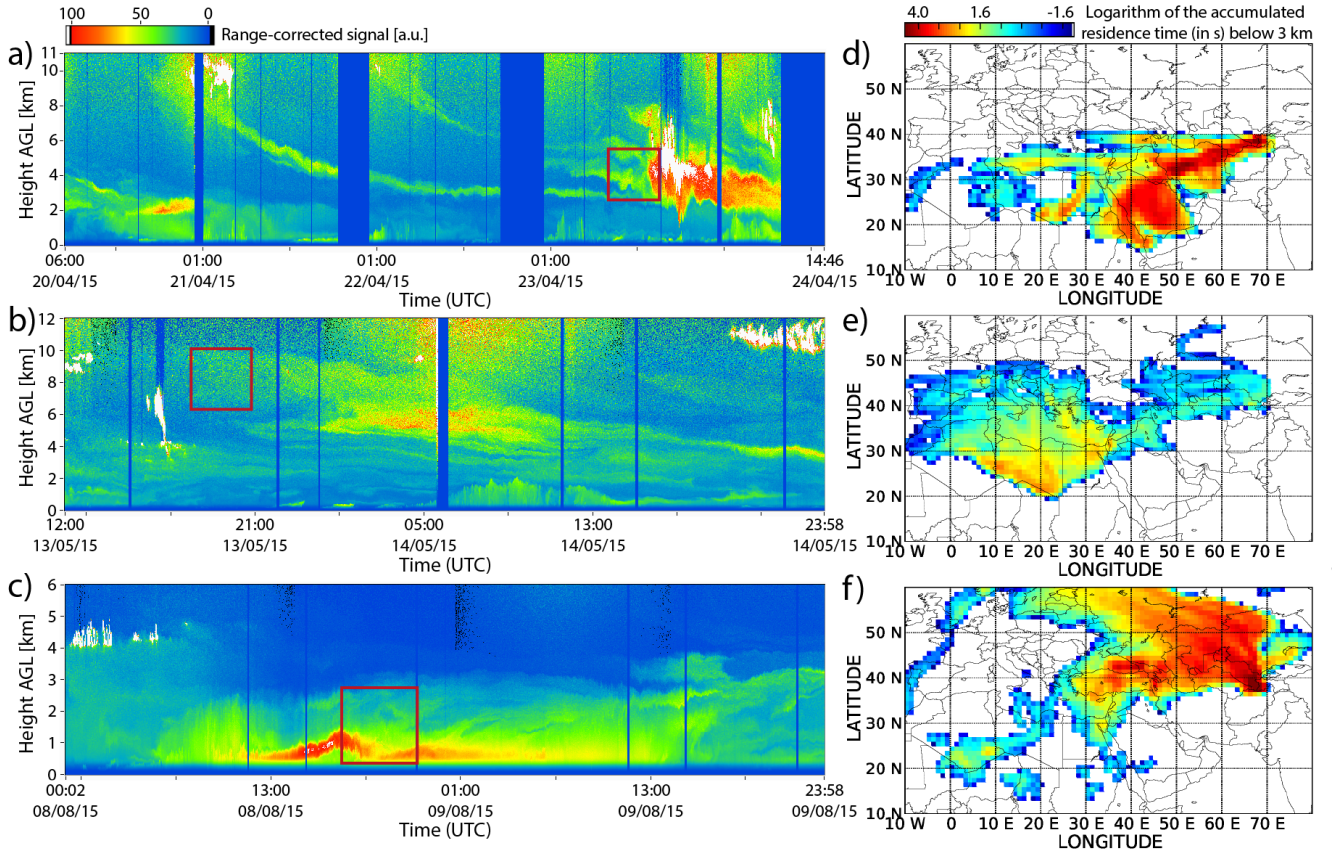


**Figure 2.** a) Polly<sup>XT</sup> with closed roof, open doors, and removed protective plates at the Poligon field site in Dushanbe. The cabinet size is approximately 1.9×1.7×0.9 m. b) Schematic sketch of the optical layout of the Polly<sup>XT</sup>. Left (transmitter part): After the second harmonic generation (SHG) and the third harmonic generation (THG), parts of the laser beam are deflected to a power meter (PM) which measures the UV component to monitor the conversion efficiency. Right (receiver part): Backscattered light is collected with a Newtonian telescope and then passed towards the receiver unit. The numbers indicate the wavelength in nm of the detection channels and c denotes the cross-polarized channels. A camera (CAM) is synchronized to the laser trigger and sees the beam at 532 nm wavelength to monitor the overlap. The polarizer mounted in front of the pinhole is a device for the absolute calibration of the depolarization measurements (details see Engelmann et al., 2016).

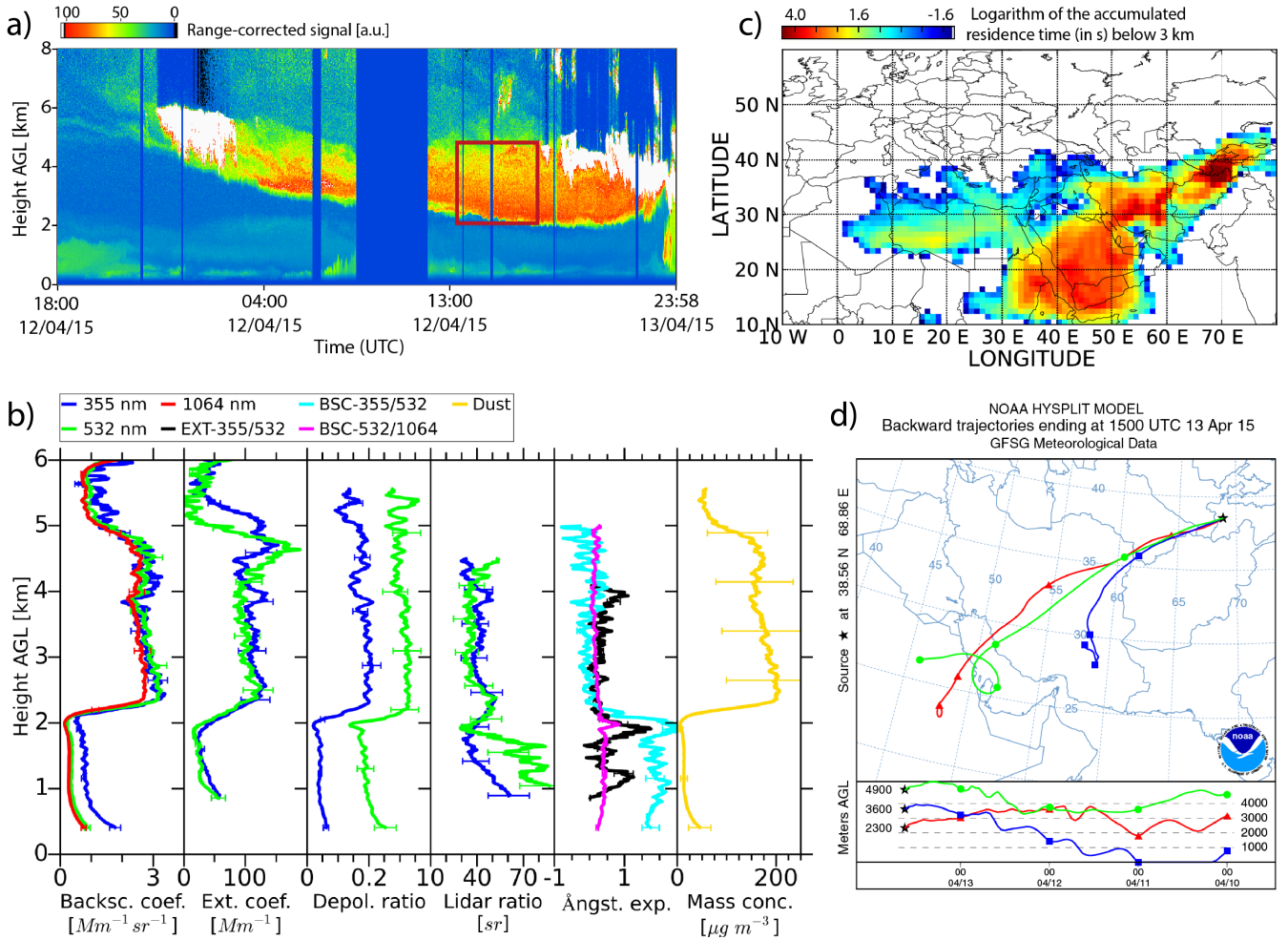




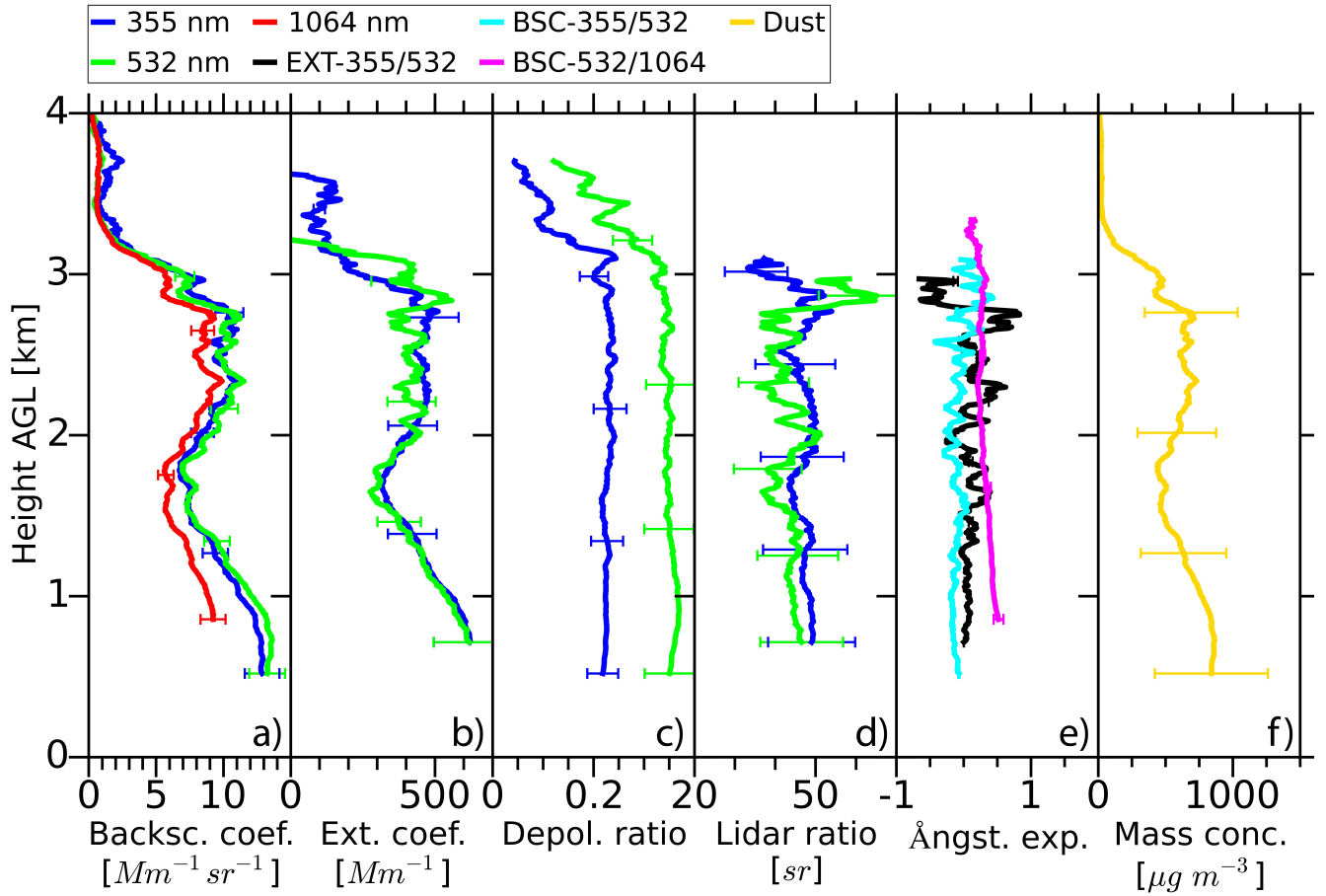
**Figure 3.** Daily means of the AOT (top) and Ångström exponent (bottom) measured in Dushanbe from March 2015 until August 2016 by sun photometer (AERONET Level 1.5, clouscreened). The red lines indicate the days of the presented example cases.



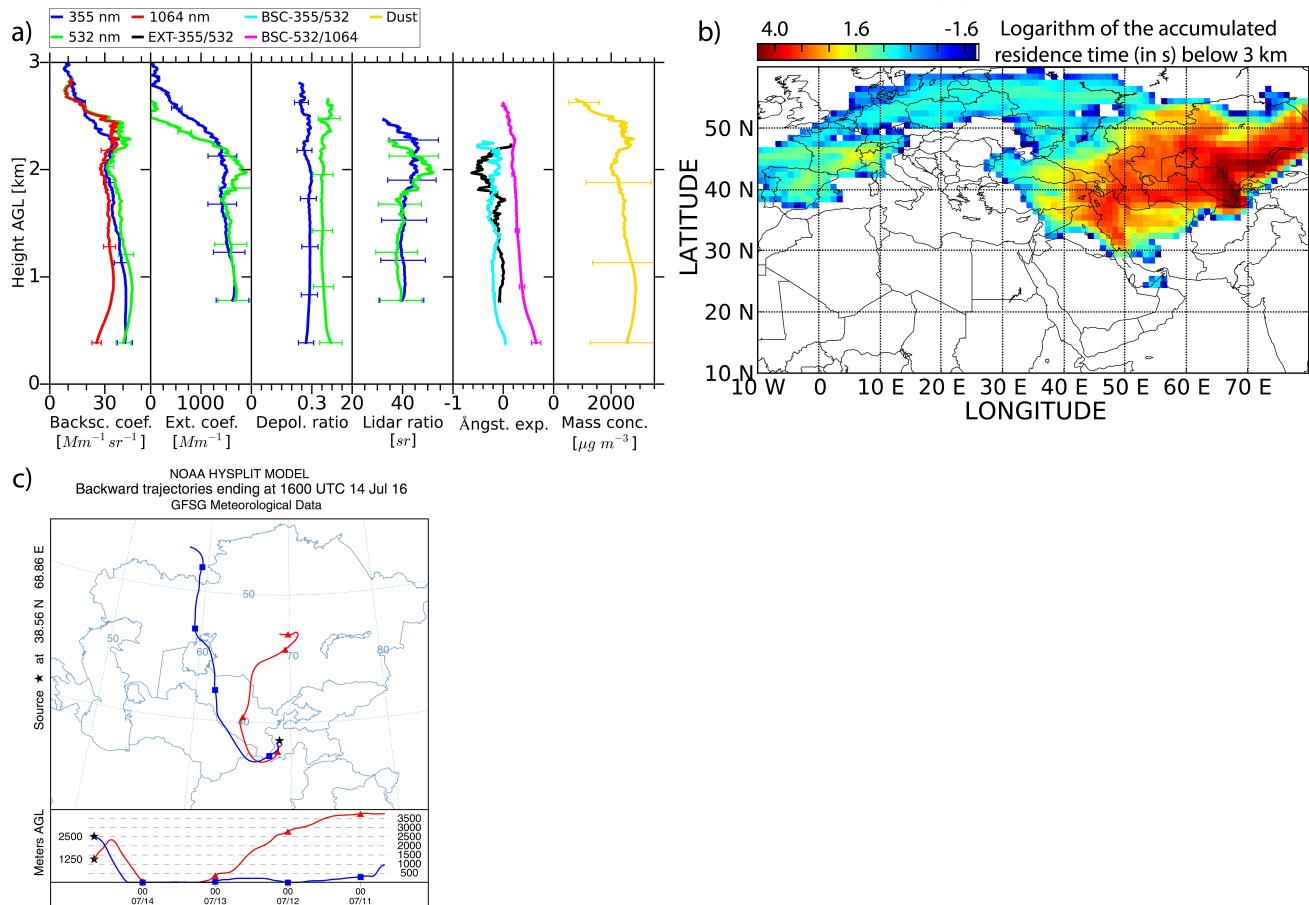
**Figure 4.** Left panels: Illustrative measurement examples of aerosol layers observed with Polly<sup>XT</sup> at Dushanbe, partly up to cirrus level (10 km height). Temporal development of the range-corrected signal (temporal resolution 30 s, vertical resolution 7.5 m) of the 1064 nm wavelength channel on a) 20 April 2015, 06:00 UTC – 24 April 2015, 14:46 UTC. b) 13 May 2015, 12:00 UTC – 14 May 2015, 23:58 UTC. c) 8 August 2015, 00:02 UTC – 9 August 2015, 23:58 UTC. Blue rectangles denote periods where no measurements were performed. Right panels: Source regions of the observed aerosol identified based on FLEXPART model runs. The first lofted layer (a) has Middle Eastern sources (d), the second lofted layer (b) has North African sources (e) and the third low laying dust (c) has Central Asian sources (f). The dark red squares in the left panels indicate the arrival height and time of the calculated backward trajectories in context of the lidar measurements. The right panels show maps of the logarithm of accumulated residence time below 3 km (in s) of 144 h FLEXPART backward trajectories arriving above Dushanbe on d) 23 April 2015, 20:30–21:30 UTC between 2.7 and 5 km, e) 13 May 2015, 17:30–18:30 UTC between 6.6 and 9.9 km height, f) 8 August 2015, 21:30–22:30 UTC between 0.5 and 3.1 km height.



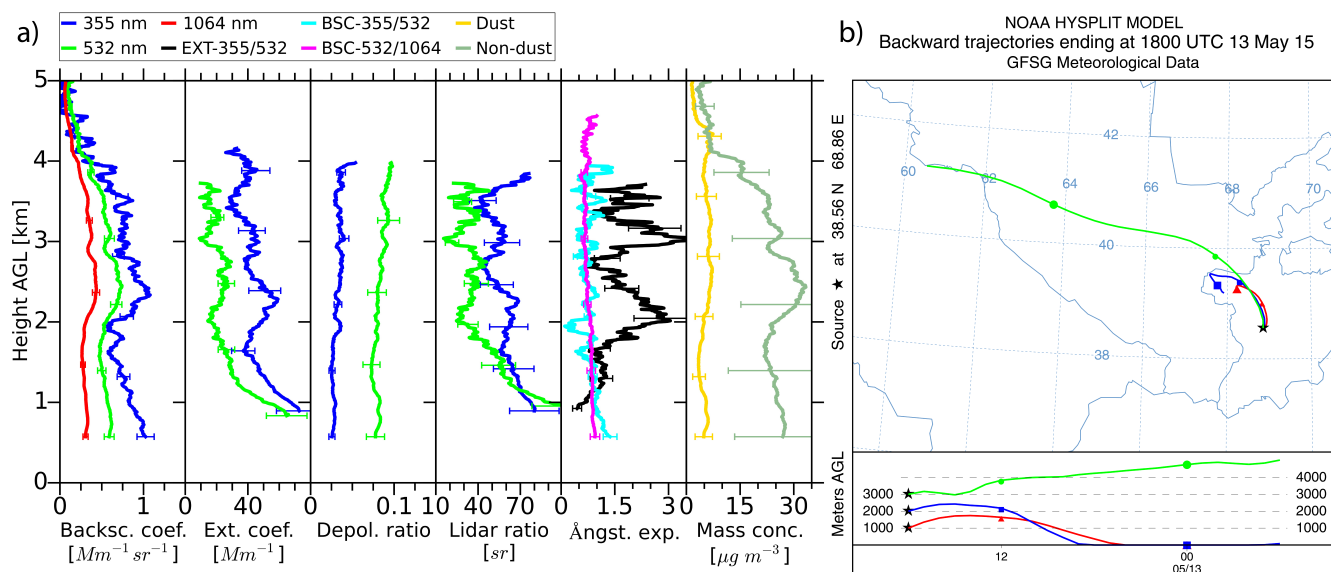
**Figure 5.** Case 1. A lofted layer of Middle Eastern dust is shown. a) Same as Fig. 4 (left panels), but on 12 April 2015, 18:00 UTC – 13 April 2015, 23:58 UTC. b) The averaged lidar profiles were measured on 13 April 2015, 15:10–16:08 UTC. Lidar signals are smoothed before calculation of the extinction-related (743 m) and backscatter-related (68 m) optical properties. Particle backscatter coefficient at 355 nm, 532 nm and 1064 nm wavelength. Particle extinction coefficient, particle linear depolarization ratio, and lidar ratio at 355 nm and 532 nm wavelength. Extinction-related Ångström exponent from 355 nm and 532 nm wavelengths (black), backscatter-related Ångström exponent from 355 nm and 532 nm wavelengths (aqua), and backscatter-related Ångström exponent from 532 nm and 1064 nm wavelengths (magenta). Dust mass concentration. c) 144 h FLEXPART backward trajectories same as in Fig. 4 (right panels), but arriving on 13 April 2015, 14:30–15:30 UTC between 2.3 and 4.9 km height. d) 96 h HYSPLIT backward trajectories arriving at Dushanbe at 18:00 UTC at 2.3, 3.6, and 4.9 km height.



**Figure 6.** Case 2. An extreme dust layer with an AOT of 1.5 (at 532 nm wavelength) and a calculated dust mass concentration of  $845 \mu\text{g m}^{-3}$  is shown. The source regions of this dust are in Central Asia (Fig. 4f). a) – f) same as Fig. 5b. The averaged lidar profiles were measured on 8 August 2015, 22:20–23:57 UTC. Lidar signals are smoothed before calculation of the extinction-related (458 m) and backscatter-related (68 m) optical properties.



**Figure 7.** Case 3. The most extreme dust event with an AOT of above 4 (at 500 nm wavelength) and a calculated dust mass concentration of  $2.8 \text{ mg m}^{-3}$  is shown. The dust source regions are in Central Asia. a) same as Fig. 5b. The averaged lidar profiles were measured on 14 July 2016, 16:00–22:00 UTC. Lidar signals are smoothed before calculation of the extinction-related (743 m) and backscatter-related (248 m) optical properties. b) 120 h FLEXPART backward trajectories same as in Fig. 4 (right panels), but arriving on 14 July 2016, 20:30–21:30 UTC between 0.5 and 2.6 km height. c) 96 h HYSPLIT backward trajectories arriving at Dushanbe at 16:00 UTC at 1.25, and 2.5 km height.



**Figure 8.** Case 4. A contrasting case with local pollution is shown. a) same as Fig. 5b. Averaged lidar profiles measured on 13 May 2015, 18:10–20:30 UTC. Lidar signals are smoothed before calculation of the extinction-related (788 m) and backscatter-related (98 m) optical properties. b) 48 h HYSPLIT backward trajectories arriving at Dushanbe at 18:00 UTC at 1, 2, and 3 km height.