



1 **Influence of NO₂ on secondary organic aerosol formation from**
2 **ozonolysis of limonene**

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26 **Abstract**

27 Limonene has a strong tendency to undergo ozonolysis to form semi-volatile and
28 low-volatility compounds that contribute to secondary organic aerosols (SOAs) both
29 outdoors and indoors. The influence of NO_2 on SOA formation from ozonolysis of
30 limonene has been evaluated using chamber experiments and the Master Chemical
31 Mechanism (MCM) coupled with a gas-particle partitioning model in this work. A
32 series of 21 indoor chamber experiments were carried out with or without NO_2 under
33 different $[\text{O}_3]_0/[\text{VOC}]_0$ ratios, and these experimental data were compared with the
34 model simulations. Agreement in SOA yields was observed between the experimental
35 observations and model simulations under varying conditions. Generally, SOA mass
36 yields are positively dependent on $[\text{O}_3]_0/[\text{VOC}]_0$ without the presence of NO_2 .
37 However, the introduction of NO_2 leads to a more complicated change in SOA yield,
38 which is shown to be related to initial $[\text{O}_3]/[\text{VOC}]$ ratios. When $[\text{O}_3]_0/[\text{VOC}]_0 > 2$, the
39 introduction of NO_2 results in an increase of SOA yield in the range of NO_2 studied in
40 this work; whereas a weak negative effect was found for SOA formation according to
41 the introduction of ~ 250 ppbv NO_2 under $[\text{O}_3]_0/[\text{VOC}]_0 < 2$ conditions. It was
42 suggested that the effect of NO_2 on SOA formation yields from limonene ozonolysis
43 is related to the competition between O_3 - and NO_3 -initiated oxidation of limonene as
44 well as the competition between RO_2+HO_2 and RO_2+NO_2 (or NO_3). Analysis of
45 aerosol chemical composition with Fourier-transform infrared spectroscopy (FTIR)
46 and modeling further confirmed that the formation of peroxy acyl nitrates (PANs) and
47 organic nitrates plays an important role in aerosol particle formation from limonene



48 ozonolysis at the presence of NO_2 . The findings here indicate that accurately
49 constraining SOA yields from NO_3 oxidation is essential to evaluate the influence of
50 NO_2 on SOA formation in some real atmosphere, for example, regions with both
51 biogenic and anthropogenic influences.

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70 **1 Introduction**

71 Globally, biogenic volatile organic compounds (BVOCs) dominate the total
72 emission of VOCs into the atmosphere, which are an order of magnitude more than
73 the emissions of anthropogenic VOCs (Guenther et al., 1995). As one class of
74 terpenoids, monoterpenes ($C_{10}H_{16}$) are estimated to account for 11% of annual
75 BVOCs emissions, where α -pinene is the most dominant monoterpene emission on
76 average, and limonene is also prevalent in some regions (Geron et al., 2000;
77 Kanakidou et al., 2005; Stroud et al., 2005). Owing to their abundance in atmosphere
78 and tendency to produce low volatile products when react with O_3 , OH, and NO_3 ,
79 monoterpenes contribute a major fraction of total atmospheric SOA (Guenther et al.,
80 1995; Griffin et al., 1999; Atkinson and Arey, 2003; Ng, et al., 2017). The ozonolysis of
81 monoterpenes, in particular, has been the focus of numerous studies for two reasons
82 (Hoffmann et al., 1997; Griffin et al., 1999; Glausius et al., 2000; Koch et al., 2000;
83 Iinuma et al., 2004; Leungsakul et al., 2005; Presto et al., 2005a, b; Jonsson et al.,
84 2006; Zhang, et al., 2006; Ng et al., 2007; Saathoff et al., 2009; Pathak et al., 2012).
85 First, reaction with ozone is a significant atmospheric sink for terpenes (Atkinson,
86 1997). Second, owing to the number or position difference of the double bond in
87 molecular bone, ozonolysis of these species will contribute differently to the
88 formation of SOA via gas-to-particle conversion (Kroll and Seinfeld 2008).

89 Limonene, one important monoterpene, is emitted in a major part from vegetation
90 (Guenther et al., 1995, 2002; Griffin et al., 1999). It is the main component of
91 essential oil extracted from citrus rind and is responsible for the strong scent of



92 oranges. Limonene is also a popular ingredient in household products, such as some
93 kinds of air fresheners and cleaners (Clausen et al., 2001; Singer et al., 2006).
94 Therefore, it is of special concern for indoor exposure considering the high production
95 of SOA and possible health risks under elevated limonene and O₃ levels (Weschler
96 and Shields, 1999; Wainman et al., 2000; Sarwar et al., 2003; Nazaroff et al., 2004;
97 Waring, 2014).

98 Limonene chemistry is especially interesting since there are two very different
99 unsaturated double bonds: endocyclic tri-substituted and exocyclic terminal. It is
100 believed that the oxidation of these double bonds, leading to the production of
101 semi-volatile and low-volatility organic compounds, results in the strong tendency of
102 limonene to form SOA. For the gas-phase reaction with ozone, it has been reported
103 that the oxidation of the endocyclic double bond ($2 \times 10^{-16} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$) occurs
104 approximately 30 times faster than for the exocyclic double bond ($7 \times 10^{-18} \text{ cm}^3$
105 $\text{molecule}^{-1} \text{ s}^{-1}$) (Shu and Atkinson, 1994; Zhang et al., 2006). Given the difference in
106 reaction rates, limonene is regarded as an excellent model system for multi-generation
107 SOA formation studies (Maksymiuk et al., 2009).

108 Although the formation of SOA mainly depends on the organic precursor and the
109 initiating oxidant (OH, O₃, and NO₃), it has also been shown that other factors, such
110 as NO_x (Nojgaard et al., 2006; Zhang et al., 2006; Draper et al., 2015), particle acidity
111 (Iinuma et al., 2004), relative humidity (RH) (Bonn et al., 2002; Jonsson et al., 2006,
112 2008a, b) and temperature (Jonsson et al., 2008b; Saathoff et al., 2009), can affect the
113 partitioning equilibrium of the produced semivolatile organics and influence the



114 reaction pathways that form SOA.

115 Recent studies of aromatic hydrocarbons, isoprene and terpene photooxidation,
116 and terpene ozonolysis have demonstrated that gas-phase organic oxidation
117 mechanisms are sensitive to NO_x (Presto et al., 2005b; Song et al., 2005; Kroll et al.,
118 2006; Ng et al., 2007a, b; Draper et al., 2015). The effect of NO_x on the formation of
119 SOA has been attributed primarily to the different reaction rates of organic peroxy
120 radicals (RO_2) with nitric oxide (NO) and HO_2 , and to the difference in volatilities of
121 the products from the two respective reaction pathways (Kroll and Seinfeld, 2008).
122 Under high- NO_x conditions, RO_2+NO dominates and results in the formation of
123 volatile organic nitrates or fragmentation products (Presto et al., 2005b). However, the
124 RO_2+HO_2 pathway, which is competitive in the absence of NO_x , leads to less volatile
125 products (e.g., peroxy acids), resulting in higher SOA yields (Johnson et al., 2005;
126 Song et al., 2007). As RO_2+NO_2 reactions have generally been considered
127 unimportant for SOA formation due to the short lifetime of peroxy nitrates (<1 s),
128 so-called “high- NO_x ” conditions typically refer to high-NO conditions. However,
129 Chan et al. recently suggested that greater attention should be paid to the effect of
130 NO_2 on SOA formation (Chan et al., 2010). In their work, SOA formation from
131 reactions between acyl peroxy radicals and NO_2 has been found higher than that from
132 RO_2+HO_2 reactions at sufficiently high NO_2 concentrations. Draper et al. investigated
133 NO_2 effect on SOA formation from ozonolysis of monoterpenes at varying
134 concentrations of NO_2 in their recent work, and they found the different NO_2 effect on
135 the observed aerosol yields for different monoterpenes at various NO_2 levels, which is



136 essentially resulted from the competition between O₃ and NO₃ to oxidize each
137 monoterpene (Draper et al., 2015). Considering the current trend of increasing
138 tropospheric NO₂ over developing areas (i.e., East Asia) (He et al., 2007) and the high
139 NO₂ concentrations in urban indoor environments (Weschler et al., 1994; WHO, 2000;
140 Dimitroulopoulou et al., 2001), re-evaluating the effect of NO₂ on SOA formation
141 from typical atmospheric VOC oxidation is relevant to predicting aerosol production
142 in some mixed region with both biogenic and anthropogenic emissions, as well as
143 particle exposure from indoor environments and polluted urban atmospheres.

144 With respect to the critical roles of the chemical reaction mechanisms for
145 atmospheric VOCs oxidations in describing SOA formation, some explicit chemical
146 mechanisms based on experimental studies have been developed and applied to
147 describe the formation of SOA from the oxidation of anthropogenic or biogenic VOCs
148 over the years (Jenkin et al., 1997; Madronich and Calvert, 1999; Carter, 2000; Griffin
149 et al., 2002; Capouet et al., 2003; Saunders et al., 2003; Aumont et al., 2005). Among
150 these near-explicit or fully explicit chemical mechanisms, MCM describes complete
151 degradation of methane and 142 non-methane VOCs (Jenkin et al., 2015). By
152 comparison to controllable smog chamber experiments, different versions of MCM
153 have been evaluated for gas phase chemistry of alkenes (Hynes et al., 2005), isoprene
154 (Pinho et al., 2005; Jenkin et al., 2015), α - and β -pinene (Saunders et al., 2003; Pinho
155 et al., 2007; Xia et al., 2008), aromatics (Bloss et al., 2005). It seems that the use of
156 this explicit mechanism may therefore provide insight into the atmospheric oxidation
157 and the formation of SOA.



158 The purpose of this work is to investigate the influence of nitrogen dioxide in
159 SOA formation following the ozonolysis of limonene. First, a systematic experimental
160 study on the effect of NO₂ on SOA formation from limonene was conducted under
161 varying NO₂ levels with different O₃/limonene ratios. SOA yields and mass
162 concentration under different conditions were investigated and aerosol chemical
163 composition was analyzed off-line with FTIR. Then the subset of MCM describing
164 limonene ozonolysis coupled with a gas-particle partitioning model was used to
165 simulate the NO₂ effect on SOA formation. Finally, the mechanism of the influence of
166 NO₂ on SOA formation from ozonolysis of limonene was thoroughly analyzed by
167 comparing experiment observations and model predictions.

168 **2 Methods**

169 **2.1 Experimental section**

170 All experiments were performed in a collapsible ~850 L Teflon FEP chamber,
171 which had been used previously for characterization of oxidation of hexenols (Gai et
172 al., 2015), SOA from photooxidation of aromatic organic hydrocarbons (Huang et al.,
173 2013), and vacuum ultraviolet photoionization mass spectra of isoprene
174 photooxidation products (Pan et al., 2012). Briefly, the chamber was irradiated with
175 UV light (300-400 nm) overnight and flushed with dry zero air until the background
176 particle number less than 10 cm⁻³ prior to each experiment. A known volume of
177 limonene was firstly injected into a 250 mL temperature-controlled glass bulb
178 vaporizer (298-598K), and then flushed into the smog chamber with purified air (4
179 L/min) for about 20 min, the concentration of which was monitored with a gas



180 chromatograph- flame ionization detection (GC-FID, Agilent 7820) during the whole
181 experiment. Typically, excessive 2-butanol was also introduced into the smog
182 chamber as OH scavenger to suppress the influence of OH radicals produced from the
183 O₃-limonene reaction. To investigate the effect of NO₂, a certain concentration of
184 mixture of NO₂/purified air was pumped into the main reactor using a 0.4 L/min
185 miniature vacuum pump. The appropriate concentration of NO₂ inside the chamber
186 was obtained by controlling the pumping time, and monitored with NO_x analyzer (TEI
187 model 42i). Ozone was produced with an ozone generator in a 1 L/min clean air
188 stream, and the oxidation of precursors inside the reactor was initiated by the
189 introducing of O₃. The O₃ concentration inside the chamber was measured in real time
190 by O₃ analyzer (TEI model 49i). Aerosol produced from the limonene/O₃ system
191 under various NO₂ concentrations was monitored on-line with a scanning mobility
192 particle sizer (SMPS, TSI, 3936, USA) consisting of a differential mobility analyzer
193 (DMA, TSI model 3080) and a condensation particle counter (CPC, TSI model 3775).
194 To investigate the chemical composition of SOA, particle products were also collected
195 with PTFE filters (Pall Life Sciences, USA) for off-line FTIR analysis in some
196 experiments. Detailed information on the chemicals and off-line FTIR analysis of
197 aerosol is provided in the Supplement.

198 **2.2 Model description**

199 To obtain insight into the effect of NO₂ on SOA formation from limonene, MCM
200 coupled with a gas-particle partitioning model was applied against smog chamber
201 experiments. The current 3.3.1 version of MCM (Jenkin et al., 2015), available online



202 (<http://mcm.leeds.ac.uk/MCM>), represents the degradation of methane and 142
203 non-methane VOCs. The subset of MCM v3.3.1 describing limonene and 2-butanol
204 oxidation, containing 2272 chemical reactions and 749 compounds, was chosen for
205 gas phase species evaluation in this work.

206 The gas-particle partitioning mechanism was first proposed by Pankow (Pankow,
207 1994a,b) and has been widely used to describe the formation of organic aerosols. In
208 principle, products from oxidation of VOCs described in MCM or the other chemical
209 mechanisms can partition between the gas phase and the particle phase according to
210 their corresponding saturation vapor pressure (p_L^0). By means of the online facility
211 UManSysProp (Topping et al., 2016), the applied p_L^0 values of each stable species
212 were predicted using the EVAPORATION method (Compernelle et al., 2016). For
213 comparison, the saturation vapor pressures in this work were calculated at 291 K or
214 298 K according to the experiment conditions. Based on the ideal gas and Raoult's law,
215 the partitioning coefficient $K_{p,i}$ ($\text{m}^3 \mu\text{g}^{-1}$), which defines the thermodynamic
216 equilibrium of a given organic species between the gas and particle phases, can be
217 determined from Eq.(1):

$$218 \quad K_{p,i} = \frac{760RT}{10^6 \bar{M}_{\text{om}} \gamma_i p_{L,i}^0} \quad (1)$$

219 where R is ideal gas constant ($8.026 \times 10^{-5} \text{ m}^3 \text{ atm mol}^{-1} \text{ K}^{-1}$), T is temperature (K),
220 \bar{M}_{om} is the average molecular weight of organic matter (g mol^{-1}), γ_i is the activity
221 coefficient of compound i and set to be 1 (assumed as ideal thin solution), and $p_{L,i}^0$ is
222 the saturation vapor pressure of i. Then A_g (gaseous concentration) and F_{om}
223 (particulate-associated concentration) can be determined according to $F_{\text{om}}/A_g =$



224 $K_p TSP$, where TSP is the mass concentration of total suspended particulate material
225 (in $\mu\text{g m}^{-3}$).

226 It is assumed that equilibrium partitioning can be achieved instantaneously;
227 therefore, no chemical processes in the aerosol phase were considered in model
228 simulations in this work. Generally, it is believed that compounds with more carbons
229 are likely to condense into the aerosol phase without consideration of oligomers that
230 exist in the aerosol phase (Seinfeld and Pandis, 1998; Xia et al., 2008). In this work,
231 385 stable species with more than three carbons, out of the 701 organic compounds,
232 were included for gas-particle partitioning simulation for oxidation of limonene.

233 **3 Results and discussion**

234 As listed in Table 1, 21 darkened chamber experiments were carried out at room
235 temperature to investigate the influence of NO_2 on SOA formation in the ozonolysis
236 of limonene, which can be classified into three scenarios according to the
237 corresponding experimental conditions. In the first, ozonolysis of limonene was
238 studied under various $[\text{O}_3]_0/[\text{VOC}]_0$ ratios without NO_2 . In the second scenario,
239 ozonolysis of limonene was investigated with ~ 250 ppbv $[\text{NO}_2]_0$ under various
240 $[\text{O}_3]_0/[\text{VOC}]_0$ ratios. In the third scenario, the influence of NO_2 on SOA yield was
241 further explored under a constant $[\text{O}_3]_0/[\text{VOC}]_0$ ratio ($[\text{O}_3]_0/[\text{VOC}]_0 \sim 2.4$). During
242 each run, temperature and RH were kept almost constant. For all these experiments,
243 no inorganic seeds were added. Despite the fact that the effect of vapor wall loss on
244 SOA formation yield may be constrained by introducing seed aerosol to promote the
245 condensation of SOA-forming vapors on it, it is found that the effect can also be



246 mitigated for the reaction system with fast oxidation rate (Nah et al., 2016). Regarding
247 the high rate coefficient of limonene ozonolysis and $[O_3]_0/[VOC]_0$ ratio >2 in most
248 experiments in this work, and even for $[O_3]_0/[VOC]_0$ ratio ≤ 2 cases, the main focus is
249 to learn the effect of NO_2 on SOA formation, it is presumed that the effect of vapor
250 wall loss does not affect the conclusions of the article. An averaged aerosol density of
251 1.3 g cm^{-3} was chosen to calculate the total mass of SOA formation from limonene
252 ozonolysis from the total measured aerosol volume by SMPS (Wirtz and
253 Martin-Reviejo, 2003; Saathoff et al., 2009).

254 The model based on the subset of MCM describing limonene ozonolysis was
255 compared to the experimental data. In general, all the initial conditions for model
256 simulations follow the experimental conditions, and the total simulation time in model
257 was equal to the duration of the smog experiments.

258 **Table 1**

259 **3.1 Gas-phase reaction**

260 Figure 1 shows the time evolution of the ozone mixing ratio, NO_2 level, the
261 precursor limonene concentrations, as well as SOA formation from experimental
262 observations and model predictions in a typical high- NO_2 and high- O_3 experiment
263 (Exp.N19). The box model studies based on the subset of the MCM for ozonolysis of
264 limonene reasonably describe most of the experimental gas-phase species
265 concentration profiles in this work. The substantial agreement between the
266 model-predicted and the experimentally observed precursor limonene shows that
267 limonene decays very rapidly following the initiation of the reaction and is depleted in



268 the first 30 minutes. The simulation also performs well in predicting changes in NO₂
269 during the reaction. Although both the predicted and observed results show roughly
270 similar temporal profiles, the simulation overpredicts O₃ by approximately 30% for all
271 data after 30min. Different from the evolution of precursor limonene, NO₂ and O₃
272 keep slowly decreasing, and SOA mass concentrations also keep increasing after the
273 first rapid reaction period, which suggests multi-step reactions leading to particle
274 formation.

275 Similar to this work, previous chamber simulation studies in other groups have
276 also shown the difficulty of predicting O₃, and under- or over-predictions appeared in
277 different cases (Chen et al., 2005; Yarwood et al., 2005; Xu et al., 2015). Although
278 the main reason resulting in the discrepancy between experimental and simulated O₃
279 concentration in this work needs further study, for simulating the formation of
280 limonene-derived SOA in this study, the performance of the model is qualitatively
281 acceptable in simulating gas-phase species (especially for precursor limonene).

282 **3.2 NO₂ effect on SOA yield**

283 **3.2.1 Chamber experiments of limonene/O₃ system in the absence or presence** 284 **of NO₂**

285 To investigate the effect of NO₂ on the limonene/O₃ system, 16 experiments
286 (N1-N16) were conducted under two typical NO₂ scenarios: low NO₂ (<1 ppbv, minor
287 NO₂ is presumed to be produced concomitantly with the generation of O₃ from
288 corona-discharge) and high NO₂ (236.7-293.4 ppbv), where the initial limonene
289 mixing ratios are approximately the same (~ 125 ppbv) under the various O₃



290 concentrations. According to the partitioning theory originally outlined by Pankow
291 (Pankow, 1994a, b) and Odum et al. (Odum et al. 1996), SOA yields are calculated
292 from the ratio of the amount of SOA formed (ΔM_0 , $\mu\text{g}/\text{m}^3$, after correction of wall
293 losses, see in supplement) to the amount of hydrocarbon consumed (ΔHC , $\mu\text{g}/\text{m}^3$),

$$294 \quad Y = \frac{\Delta M_0}{\Delta\text{HC}} \quad (1)$$

295 As listed in Table 1 and shown in Figure 2, the results show two striking features.
296 In the absence of NO_2 (N1-N7), SOA yields (black squares in Figure 2) increase
297 smoothly with increasing initial O_3 concentration at almost the same $[\text{VOC}]_0$, which
298 indicates the contribution of the different generation products according to the extents
299 of limonene ozonolysis to SOA formation (Ng et al., 2006). In the presence of NO_2 ,
300 however, change in SOA yields becomes more complicated (N8-N16, red circles in
301 Figure 2). It is shown that at $[\text{O}_3]_0/[\text{VOC}]_0 < 2$ conditions (N8-N12), the introduction
302 of NO_2 results in a weak negative effect on the SOA yields compared to that without
303 NO_2 . While at $[\text{O}_3]_0/[\text{VOC}]_0 > 2$ conditions (N13-N16), SOA yields increase
304 substantially, far exceeding those without NO_2 at a similar mixing ratio of $[\text{O}_3]_0$ and
305 $[\text{limonene}]_0$ following the introduction of NO_2 in the reaction system. And this
306 increase cannot be attributed simply to the increase of initial O_3 concentration in
307 comparison to the experimental results of N4-N7. It seems that NO_2 , entangled with
308 the initial O_3 concentration, participates in some key reaction pathways of
309 limonene/ O_3 system influencing SOA formation.

310 **Figure 2**

311 **3.2.2 NO_2 experiments under a constant $[\text{O}_3]_0/[\text{VOC}]_0$ ratio with excess O_3**



312 To further investigate the effect of NO₂ on SOA formation in the limonene/O₃
313 reaction system, the third group of experiments (N17-N21) were conducted by adding
314 different concentrations of NO₂ to limonene/O₃ at the beginning of the experiments,
315 while the [O₃]₀/[VOC]₀ ratio was kept roughly constant ([O₃]₀/[VOC]₀ =2.5). In these
316 experiments, limonene completely reacted (~122 ppbv) due to high O₃ concentration
317 (~300 ppbv).

318 As shown in Figure 3, given the similar relative humidity (14.4% ± 0.6%),
319 temperature (25.6 ± 0.6 °C), initial limonene concentration (122.2 ± 3.2 ppbv), and
320 initial ozone concentration (~300 ppbv), aerosol formation increases as NO₂ added to
321 the system increases. At relatively low NO₂ concentrations (<200 ppbv), the SOA
322 yield and mass concentration were found to increase rapidly with increasing NO₂;
323 whereas the rate of increase tends to slow as the initial NO₂ concentration added
324 exceeds 300 ppbv. Given the constant [O₃]₀/[VOC]₀ ratio with excess O₃, both particle
325 concentration and particle size increase with the NO₂ level progressively, e.g., the
326 mode diameter increases from 95 to 107 nm and the total concentration increases from
327 1.77×10⁵ to 2.17×10⁵ particles/cm³ by adding 298.5 ppbv NO₂ (Figure S2 in
328 supplement).

329 **Figure 3**

330 SOA formation from limonene ozonolysis has also been examined by Zhang et al.
331 under different NO_x (NO + NO₂) levels and O₃ levels (Zhang et al., 2006). It was
332 found that heterogeneous oxidation of the first-generation products by ozone uptake
333 on unsaturated SOA particles are rate-limiting at low NO_x, while at high NO_x,



334 gas-phase ozonolysis of the terminal double bond dominates. As shown in Figure S3,
335 the SOA yields under high ratios of ozone-to-limonene ($\gg 10$) without NO_x in their
336 work is far higher than that observed in this work, which indicates the heterogeneous
337 oxidation of the exo bond follow immediately the initial oxidation of endo double
338 bond. Donahue et al. have explored SOA formation from ozonolysis of limona ketone,
339 a potentially important intermediate in limonene ozonolysis (Donahue et al., 2007).
340 And it was found that SOA formation from limona ketone is similar to that from
341 a-pinene, but a factor of four less intense than from limonene at low
342 ozone-to-limonene ratio (<1) (see the lower dashed line in Figure S3). It seems that
343 SOA yields from limonene + ozone of this work fall into the range between the two
344 aforementioned work, indicting the dependence of SOA formation on the level of
345 multigenerational oxidation of limonene that can occur. At the presence of NO₂, the
346 trend of SOA yields is more complex due to differences in the volatility of the entire
347 product suite formed from different reaction passways. These sensitivities to
348 ozone-to-limonene ratio and NO_x suggest that great care should be taken in
349 extrapolating laboratory results to real atmosphere.

350 **3.2.3 Comparison of model-predicted SOA yields with experimentally observed** 351 **yields**

352 To explore the NO₂ effect on SOA formation during the ozonolysis of limonene,
353 the simulated SOA yields were explored under different NO₂ conditions, which were
354 further compared with the experimental data. Figure 4 shows the dependence of SOA
355 yield (both observed and predicted) on NO₂ level, as well as O₃ level, for oxidation of



356 limonene. Reasonable agreement is shown between the trend of the predicted SOA
357 yields and the experimentally observed results according to the change of NO_2 , which
358 indicates that the MCM /gas-aerosol partitioning model adequately captures the effect
359 of NO_2 on SOA yield for the limonene/ O_3 reaction system. For example, the
360 model-predicted SOA yields tend to increase explicitly with increasing initial O_3
361 concentrations with (Exp N8-N16 series) or without NO_2 (Exp N1-N7 series),
362 consistent with the experimental results discussed in 3.2.1. As for the constant
363 $[\text{O}_3]_0/[\text{VOC}]_0$ ratio with excess O_3 , the predicted SOA yields, increasing with
364 increasing $[\text{NO}_2]_0$, are roughly equal to the observed SOA yields except at 611 ppbv
365 $[\text{NO}_2]_0$, where the predicted SOA yield is less than the observed value. It is
366 noteworthy that the two different temperatures were used for model simulation in
367 Figure 4 (N1-N16:291K; N17-N21:298K) according to the experimental conditions.
368 As the slope of the fit line is 1.4, it means that the predicted results are higher than the
369 observed ones. It also can be found that the predicted yields at the lower temperature
370 are far higher than that at higher temperature (see comparison in Figure S4 in
371 supplement). SOA formation from ozonolysis of limonene has been investigated as a
372 function of temperature in the range of 253–313 K, and negative dependence of SOA
373 yields on temperature has been found (Saathoff et al., 2009). Based on their work and
374 the corresponding organic aerosol mass in this work, 15% increase of SOA yields at
375 the utmost can be attributed to the 7 K temperature difference between Exp.N1-N16
376 and N17-N21. However, the predicted yields at 291K(N1-N16) is 1.5 to 2 times of
377 that at 298K(N1-N16) (Figure S4 in supplement), it is presumed that this big



378 difference maybe result from the imprecise evaluation for the saturation vapor
379 pressures of the different kind of products of limonene ozonolysis.

380 **Figure 4**

381 The sensitivity of the SOA yield to NO_2 and O_3 was explored further based on 861
382 simulations, with the concentration of O_3 ranging from 0 to 800 ppb and NO_2 ranging
383 from 0 to 400 ppb, while the initial concentration of limonene was set as 125 ppb,
384 close to the typical concentration of limonene adopted in the chamber experiments in
385 this work (See Table 1). Figure S5a in supplement shows the simulated SOA yield as
386 a function of O_3 and NO_2 level at 291K. Most SOA yields here refer to the yields
387 when >90% of initial limonene loading was consumed. Generally, higher yields can
388 be observed for higher O_3 level or NO_2 level as aforementioned. As shown in Figure
389 S5b, at the two typical NO_2 concentrations (0.1 and 0.26 ppm similar to the
390 experimental conditions), the simulated SOA yields show a positive correlation with
391 the initial O_3 concentrations no matter the concentration of NO_2 , which is consistent
392 with the experimental trend. However, the relationship between the SOA yields and
393 the initial NO_2 levels are more complicated and entangled with O_3 concentration. As
394 shown in Figure S5c, in the presence of 100 ppb O_3 , the simulated SOA yield first
395 decreases with increasing NO_2 at low concentration (below 50 ppb). This may
396 qualitatively explain the experimental observation that SOA yields in the presence of
397 NO_2 are smaller than those in the absence of NO_2 under $[\text{O}_3]_0/[\text{VOC}]_0 < 2$ conditions,
398 as shown in Figure 2. Then the simulated SOA yield increases with increasing NO_2 at
399 the moderate region (ranging from 0.05 to 0.3 ppmv), and finally decreases again with



400 increasing NO_2 when concentrations are even higher. In the presence of 300 ppbv O_3 ,
401 simulated SOA yields show an increasing trend with increasing NO_2 until
402 approximately 0.3 ppmv, and a decreasing trend with increasing NO_2 after this point.
403 For comparison, the experimental data of N17-N20 and the model predicted results at
404 298K, where $[\text{O}_3]_0$ was equal to the simulation conditions (~ 300 ppbv), are also
405 shown in Figure S4c. It is obvious to see the difference resulting from the predictions
406 in different temperatures. In the presence of even higher O_3 (0.7 ppmv here), the
407 simulated SOA yield first increases and then levels off with increasing NO_2 .

408 3.3 NO_2 effect on SOA chemical components

409 3.3.1 FTIR analysis of SOA particles formation in the limonene/ O_3 system

410 Traditionally, the effect of NO_x on SOA formation was believed to result from the
411 competition between the reaction of RO_2+HO_2 and that of RO_2+NO . The RO_2+HO_2
412 pathway, which dominates at low NO_x condition, tends to produce less volatile
413 organic hydroperoxides (ROOHs) and peroxy acids, leading to the formation of SOA.
414 Under high-NO conditions, the reaction of RO_2 with NO will result in the production
415 of volatile organic nitrates. However, this reaction pathway can be ruled out at high
416 O_3 conditions. Although RO_2+NO_2 reactions have been considered less important
417 before for SOA formation, it was confirmed recently that the formation of peroxy acyl
418 nitrates (PANs, $\text{RC}(\text{O})\text{OONO}_2$) contribute certainly to SOA formation from the
419 oxidation of sesquiterpenes (Ng et al., 2007) and aldehydes (Chan et al., 2010) at high
420 NO_2 levels. While in the recent literature, reactions of NO_3 with BVOCs have
421 received increased attention and been recognized as a potential source of SOA from



422 ozonolysis of BVOCs at high NO₂ concentration (Ng et al., 2017).

423 To learn the effect of NO₂ on the SOA chemical component, the filtered
424 particle-phase products from the chamber limonene/O₃ reaction were analyzed by
425 FTIR. Here, the FTIR spectra of particles separately collected from high NO₂
426 experiment (~660 ppbv) and low NO₂ experiment (<1 ppbv), as well as the spectra for
427 a blank PTFE filter, are shown in Figure 5. Compared with the base absorption
428 spectrum, the samples from oxidation of limonene feature absorption at 1720 cm⁻¹,
429 which indicates the production of compounds with a carbonyl group, such as
430 aldehydes, ketones, or acids (Leungsakul et al., 2005; Presto et al., 2005b). Disparate
431 features are evident in the IR spectra of high-NO₂ experiment as compared with
432 no-NO₂ experiment. Following previous experimental works (Presto et al., 2005b;
433 Palen et al. 1992; Bruns et al., 2010) and field work (Mylonas et al. 1991), these
434 features are assigned as characteristic organic nitrate IR peaks, according to NO
435 symmetric stretch (860 cm⁻¹), NO₂ symmetric stretch (1280 cm⁻¹) and NO₂
436 asymmetric stretch (1630-1640 cm⁻¹) (Socrates et al. 2001).

437 **Figure 5**

438 FTIR analysis of SOA particles in this work has identified the contribution of
439 organic nitrates to SOA formation in limonene/O₃ system in the presence of high NO₂.
440 Actually, the contribution of low-volatility organic nitrate to total particle mass has
441 also been observed in the SOA formation from limonene ozonolysis (Zhang et al.,
442 2006) or photooxidation of sesquiterpenes (such as longifolene and aromadendrene)
443 (Ng et al., 2007) at the presence of NO_x. Recently, organic nitrate in SOA formation



444 from ozonolysis of limonene has been evaluated at varying concentration of NO₂, and
445 the organic nitrate yield from NO₃ + limonene is found to be about 30% (Fry et al.,
446 2011; Draper et al., 2015). Regarding to the similar reaction conditions to the work in
447 Fry et al. and Draper et al., organic nitrate observed here is suggested to be resulted
448 from direct NO₃ oxidation of limonene or indirect NO₂/NO₃ reaction with the
449 reactive intermediate compounds following ozonolysis of limonene.

450 **3.3.2 Model results of the NO₂ effect on total SOA**

451 As mentioned, our FTIR analysis of SOA particles has confirmed that compounds
452 with carbonyl groups contribute to the particle phase regardless of the presence of
453 NO₂, while organic nitrates contribute to the particle phase only in the presence of
454 NO₂. To evaluate the effect of NO₂ on aerosol formation, simulation work was also
455 carried out based on MCM model. The total mass (unit: μg m⁻³) of the four groups of
456 main products (ROOHs, acids, PANs, and nitrates (PANs are excluded)) in both gas
457 and aerosol phases with the five typical different initial mixing ratios of NO₂ are
458 predicted at the assumed conditions of 125 ppbv [limonene]₀, 300 ppbv [O₃]₀ and
459 291K.

460 As shown in Figure 6, both the total mass (gas phase + aerosol phase) of ROOHs
461 and of acids decrease with increasing NO₂, which is interpreted with the recognized
462 mechanism of RO₂+HO₂ reaction (Johnson et al., 2005). However, their relative
463 proportions to gas phase and aerosol phase show different trends following changes in
464 NO₂ concentration. It is obvious that in gas phase, the relative proportion of ROOHs
465 decreases very fast with increasing NO₂; while in aerosol phase, ROOHs increases



466 slightly when NO_2 increases from 0.0 ppmv to 0.1 ppmv; and then decreases slowly
467 with increasing NO_2 . Even in high NO_2 conditions, ROOHs still mostly contribute to
468 organic aerosols. While the fraction of acids decrease monotonically both in gas phase
469 and aerosol phase. Given its high volatility, it is obvious that PANs increase greatly in
470 the gas phase with increasing NO_2 . Considering that the increase of PANs in the gas
471 phase in turn increases its partitioning into the particle phase, it is reasonable to
472 suggest that PANs may also contribute its proportion to organic aerosol. In this
473 simulation work, PANs show the same trends both in the gas phase and aerosol, first
474 increasing with increasing NO_2 from 0.1 ppmv to 0.3 ppmv, and then decreasing
475 slowly with increasing NO_2 . As shown in Figure 6b, the particle phase PANs
476 contribute the second proportion to total particle compounds. Organic nitrates increase
477 both in gas and particle amount with the increase of NO_2 , too. And the particle phase
478 nitrates, which are also believed to result from increased partitioning, contribute the
479 third proportion to total particle compounds.

480 Regarding the model conditions, for example, 125 ppbv [limonene]₀ and 300
481 ppbv [O_3]₀, the contributions of ROOHs, acids, PANs and organic nitrates to particle
482 phase according to the changing of NO_2 concentration appear to result in a
483 corresponding changing trend of SOA yield with changing NO_2 (see Figure S5c). The
484 simulation result here also gives a semi-qualitative explanation for the FTIR analysis
485 of SOA particles.

486 **Figure 6**

487 **3.3.3 Model Investigation of SOA Chemical Composition by Individual Product**

488 **Species**

489 Although 385 condensable compounds have been examined in this simulation,
490 most contribute negligible fractions to total aerosol mass. By assessing the
491 corresponding aerosol mass fractions at 5 different NO₂ scenarios, the 20 dominant
492 condensable compounds were identified. As summarized in Table 2, most of these
493 compounds have carbon numbers larger than seven, with the two exception of
494 C628OOH (a six-carbon compound) and C735OOH (a seven-carbon compound),
495 which confirmed that organic compounds with more than six carbons have major
496 potential to form SOA (Seinfeld and Pandis, 1998). As shown in Table 2, 10 of the 20
497 dominant condensable compounds are ROOHs, and the rest include 4 acids, 3 PANs
498 and 3 nitrates. Their relative aerosol mass fractions to total SOA mass at different
499 initial concentrations of NO₂ are shown in Figure 7. Across all simulations at five
500 different NO₂ scenarios, the selected 20 dominant condensable compounds cover
501 93-96% of the total SOA mass. The simulation results indicate that most of the
502 individual compounds follow the same pattern for the grouped compounds as
503 illustrated in Figure 6b. Most of the ROOHs and acids contribute more to the total
504 SOA mass at low NO₂ conditions, while their relative importance significantly
505 decreases with increasing NO₂. On the contrary, the relative contributions of the most
506 PANs and nitrates to total SOA increase with increasing NO₂. It is interesting to note
507 some exceptions. Both NLIMOOH(C₁₀H₁₇NO₅) and NLIMALOOH (C₁₀H₁₇NO₇)
508 are ROOHs, and their mass fractions increase as the NO₂ concentration increases. In
509 particular, NLIMOOH has an aerosol mass fraction above 50% relative to the total



510 SOA mass when NO_2 concentration is higher than 0.2 ppmv. Although NLIMOOH
511 and NLIMALOOH are ROOHs, the two compounds also contain an $-\text{ONO}_2$ group in
512 the molecular blocks (see Table S1). Detailed reaction pathway analysis (in the
513 following 3.4 section) show that their production are related to NO_3 reactions, so it is
514 reasonable to find their mass fractions increasing with increasing NO_2 . For
515 C823PAN(C₉H₁₃NO₇), the result is opposite, its mass fraction decreases rather than
516 increases with increasing NO_2 . Although it is traditionally attributed to PANs,
517 C823PAN also has a carboxyl group (C(O)OH). The reaction pathway analysis (in the
518 following 3.4 section) show its production are related to RO_2+NO_2 reactions, which
519 maybe explain why its trend changes with changing NO_2 .

520 **Figure 7**

521 3.4. Reaction pathway analysis

522 To understand the relationship between SOA yield and NO_2 level, as well as the
523 individual aerosol compositions and NO_2 level, the mechanism of oxidation of
524 limonene by O_3 in the presence of NO_2 should be considered comprehensively. As
525 shown in Scheme 1, NO_2 can affect SOA formation from ozonolysis of limonene in
526 two different ways: it can form NO_3 firstly (Atkinson, 1997) that can compete with O_3
527 to oxidize limonene directly, or react with RO_2 radicals or other products and reaction
528 intermediates as NO_2 or NO_3 .

529 **Scheme 1**

530 Although a few of work reported that the nucleation potential of O_3 is much
531 higher than that of NO_3 (Bonn et al., 2002; Nojgaard et al., 2006), there are more



532 laboratory and chamber studies suggested that NO_3 - BVOC chemistry is a potentially
533 efficient SOA formation mechanism, and some SOA yield from NO_3 oxidation of
534 terpenes are even greater than that for OH or O_3 oxidation (Ng et al., 2017). The
535 model study here confirmed that the reaction of NO_3 with limonene at elevated NO_2
536 and O_3 plays an important role in the formation of SOA. Firstly, the initial
537 competition between O_3 and NO_3 oxidation directly results in distinct decrease of
538 ROOHs and acids with the increasing NO_2 . Secondly, NO_3 - initiated oxidation of
539 limonene itself can also produce some condensable compounds contributing to SOA
540 formation. For example, both NLIMOOH, the largest contributor to the total SOA
541 mass, and the other dominant condensable compound LIMBNO₃ are products from
542 NO_3 +limonene reaction passage (Figure 7). In fact, large amounts of yet unidentified
543 organic nitrates (approximately 30%) have also been found experimentally as the
544 major reaction product of limonene oxidation by NO_3 (Spittler et al., 2006; Fry et al.,
545 2011), which substantiates the observation in this work.

546 NO_2 effect on SOA formation is embodied not only in the initial competition
547 between O_3 and NO_3 oxidation, but also in participation in the production of PANs
548 and nitrates following ozonolysis of limonene. The first step of limonene ozonolysis is
549 dominated by cycloaddition of O_3 to the endocyclic double bond (Leungsakul et al.,
550 2005; Pan et al., 2009). Two highly reactive Criegee intermediates (CI^* , named
551 LIMOOA and LIMOOB here) are produced following the breakage of limonene's
552 cyclohexene ring. However, subsequent reactions of the CIs involve various reactions
553 with RO_2 or HO_2 , as well as NO_2 or NO_3 (see details in Scheme S1 and Scheme S2 in



554 supplement), which will affect the formation of SOA. For example, production of
555 C822PAN (a C9 PAN) and C822CO2H (a C9 acid) is involved in the competition
556 between the reactions of C822CO3 (a C9 peroxy radical as an intermediate product of
557 LIMOOA) with RO₂/HO₂ and with NO₂. While LIMONIC (abbreviation for limonic
558 acid, See Table 2), one main aerosol component identified experimentally (Walser et
559 al., 2008) and theoretically (Pathak et al., 2012), and C823PAN are also related to the
560 competition between the reactions of C823CO3 (another C9 peroxy radical with a
561 carboxyl) with RO₂/HO₂ and with NO₂. In the following reactions related to LIMOOB,
562 LIMBOO (stabilized form of LIMOOB) tends to react with H₂O in the absence of
563 NO₂, leading to the formation of limononic acid (abbreviated as LIMONONIC, See
564 Table S1), one of the main identified aerosol-phase acids (Pathak et al., 2012; Walser
565 et al., 2008). While in the presence of NO₂, LIMBOO will react with NO₂ to produce
566 limonaldehyde firstly (abbreviated as LIMAL here), which is followed with
567 multiple reactions leading to the formation of a series of condensable compounds. For
568 example, reaction of LIMAL with RO₂ will also form LIMONONIC. Whereas
569 multi-step reactions of LIMAL with NO₃ will lead to the generation of NLIMALOH
570 and C817PAN, two dominant contributor to particle mass as listed in Table 2 and
571 shown in Figure 7. Considering the competitions between the different reaction
572 pathways related to NO₂(NO₃) and RO₂(HO₂), it is reasonable to see the different
573 profile of the various reaction products contributing to aerosol mass according to the
574 different NO₂ levels as discussed in part 3.3.3 and shown in Figure 7.

575 **4 Conclusions**



576 This work has systematically investigated the effect of NO₂ on SOA yields and
577 composition from ozonolysis of limonene. Although the concentrations for the
578 gaseous compounds of interest (limonene, NO₂ and O₃) in real atmosphere is not
579 exactly same to precursor concentrations used in this work, and the chemistry in real
580 atmosphere is also more complicated than that in smog chamber and box model
581 studied here, thus absolute observed aerosol yields are not atmospherically relevant
582 directly. That the ratios of [NO₂]₀ : [limonene]₀ are approximately 2 : 1 (Exp.N8-N16)
583 and [NO₂]₀ : [O₃]₀ range from 3.7 : 1 to 0.4 : 1 (Exp. N8-N21 except) are reasonable
584 for relatively polluted while with high BVOCs emission sites in the real atmosphere
585 or some indoor conditions.

586 Both experimental observation and model prediction have demonstrated that the
587 effect of NO₂ on the formation of SOA from a limonene/O₃ reaction mixture is more
588 complicated than previously acknowledged. Analysis of the reaction pathway and the
589 related aerosol chemical composition suggests that NO₂ is not only related to the
590 initial competition between O₃ and NO₃ oxidation of limonene, but also the
591 competition between RO₂+HO₂ and RO₂+NO₂(or NO₃) following the ozonolysis of
592 limonene. The important role of peroxy acyl nitrates (PANs) and organic nitrates, in
593 addition to ROOHs and acids, have been identified in aerosol particle formation from
594 limonene ozonolysis with the presence of NO₂. And fractions relative to the total SOA
595 mass from PANs, organic nitrates, ROOHs and acids have been found sensitive to the
596 change of the [O₃]₀/[VOC]₀ ratios and [NO₂]₀ : [VOC]₀ ratios.

597 Actually, understanding the effects of anthropogenic emissions on aerosol



598 formation from biogenic emissions and to answer to what extent can biogenic SOA be
599 controlled is a critical area whenas a challenge of current research (Ng et al., 2017). In
600 a series recent field campaigns carried in the U.S. (Rollins et al., 2012; Setyan et al.,
601 2012; Fry et al., 2013; Xu et al., 2015a, 2015b) and in Europe (Kulmala et al., 2011;
602 Crippa et al., 2013; Hao et al., 2014; Kiendler-Scharr et al., 2016), organic nitrates are
603 found ubiquitous in SOA formed from the reaction of nitrate radicals (NO_3) and
604 BVOCs. For example, Xu et al. studied the effect of anthropogenic NO_x on biogenic
605 aerosol formation in the southeastern US, and found that monoterpenes+ NO_3
606 chemistry accounts for 50% of total nighttime OA production (Xu et al., 2015a), and
607 organic nitrates can contribute 63-100% of total nitrates in the southeastern US (Xu et
608 al., 2015b). As in a recent field campaign carried in Europe, organic nitrates were
609 found to contribute one-third of the total nitrate mass in a boreal forestland–urban
610 mixed region, and the formation of organic nitrates was also suggested as the outcome
611 of BVOCs undergoing nighttime chemistry with NO_3 (Hao et al., 2014), which
612 confirmed the findings in other location in Europe (Kiendler-Scharr et al., 2016).
613 Considering their measurement site is influenced by both biogenic and anthropogenic
614 sources, similar to the situations researched in this work, their observation confirmed
615 the suggestions in this work that the effect of NO_2 on SOA yields from
616 atmospherically relevant BVOCs, especially under the complex atmosphere with high
617 O_3 and NO_2 , should be evaluated more systematically than before, and more work are
618 needed in order to more accurately treat aerosol formation for regions with substantial
619 anthropogenic-biogenic interactions and predict air quality and climate.



620 **Acknowledgements.** The authors gratefully acknowledge support from the National
621 Natural Science Foundation of China (41575126, U1532143, 91544228, 41575125,
622 41375147, 41575118, U1232209).

623 **Supporting Information**

624 Chemicals, experimental details for off-line FTIR analysis, correction for SOA wall
625 loss, Figures S1–S6 and Scheme S1-S2 are available in the Supporting Information.

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1016 **Figure Captions:**

1017 **Figure 1.** Observed (points) and predicted (lines) concentration profiles of limonene
1018 (purple squares), O₃ (blue filled triangles), NO₂ (red dots) and SOA
1019 (diamonds-raw; filled squares-corrected) versus reaction time for Exp. N19.

1020 **Figure 2.** Aerosol yield for the limonene/O₃/NO₂ system for the initial terpene mixing
1021 ratios of ~ 125 ppb with NO₂ (~250 ppbv) or without NO₂. Each data point
1022 represents an individual experiment.

1023 **Figure 3.** SOA growth as a function of initial NO₂ concentration under the constant
1024 [O₃]₀/[VOC]₀ ratio ([O₃]₀/[VOC]₀ ratio~2.5).

1025 **Figure 4.** Comparison of the observed SOA yield and the predicted SOA yield for all
1026 the smog chamber experiments. The sizes of the data points represent the
1027 corresponding initial O₃ levels of the simulated experiments, and the color scale
1028 represents corresponding initial NO₂ levels. The line is linear fit to illustrate the
1029 relativity between the predicted SOA yield and the observed SOA yield.
1030 (Number 1-21 above the data points corresponds to Experiment N1-N21 shown
1031 in Table 1)

1032 **Figure 5.** FTIR spectra of the particle products collected with PTFE filters from a
1033 high NO₂ experiment (upper panel), a low NO₂ experiment (medium panel), and
1034 a blank experiment (lower panel).

1035 **Figure 6.** Changes of ROOHs, acids, PANs, and nitrates with different initial mixing
1036 ratio of the NO₂ (A) in gas phase, (B) in aerosol phase, and (C) in both phase.

1037 **Figure 7.** Aerosol mass fraction of the 20 dominant condensable compounds at
1038 different initial concentration of NO₂.



1039 **Scheme 1.** Major reaction paths in the initial ozonolysis and NO₃ oxidation of
1040 limonene.(Species in solid boxes are condensable compounds which can be
1041 partition to the organic aerosol phase. Species in solid boxes with shadow are
1042 those contribute dominantly to SOA in this simulation as listed in Table S1.
1043 Species in dashed box refer to the reactive radicals which generally will react
1044 further. Broken arrows indicate there are multi-step reaction not shown, which
1045 result in further degradation of the related species.)

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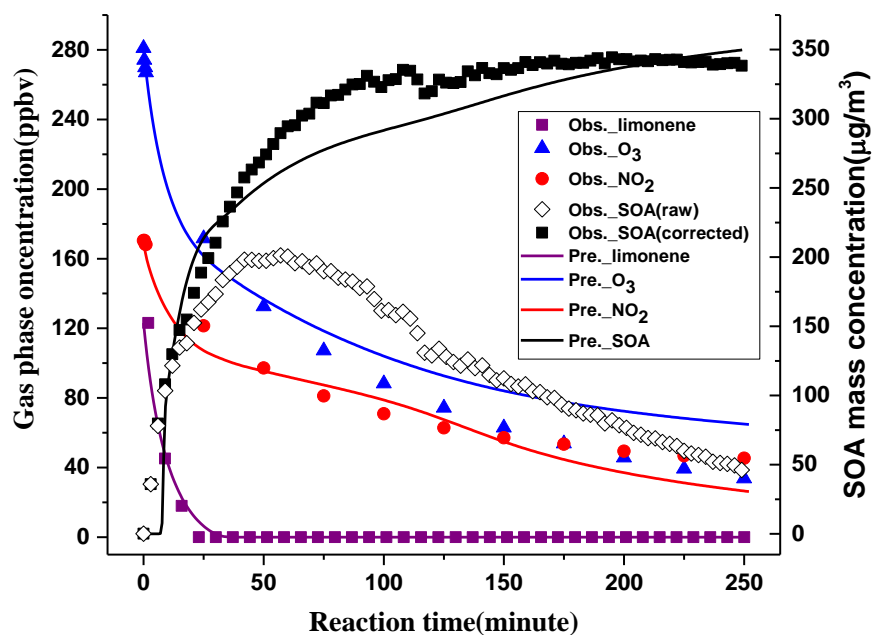
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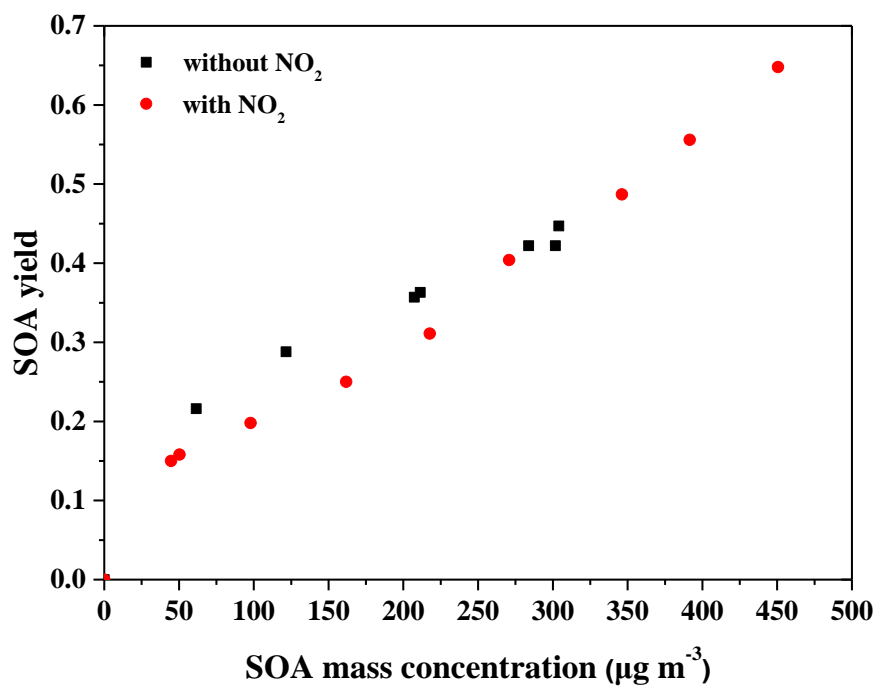
Figure 1. Hu *et al*



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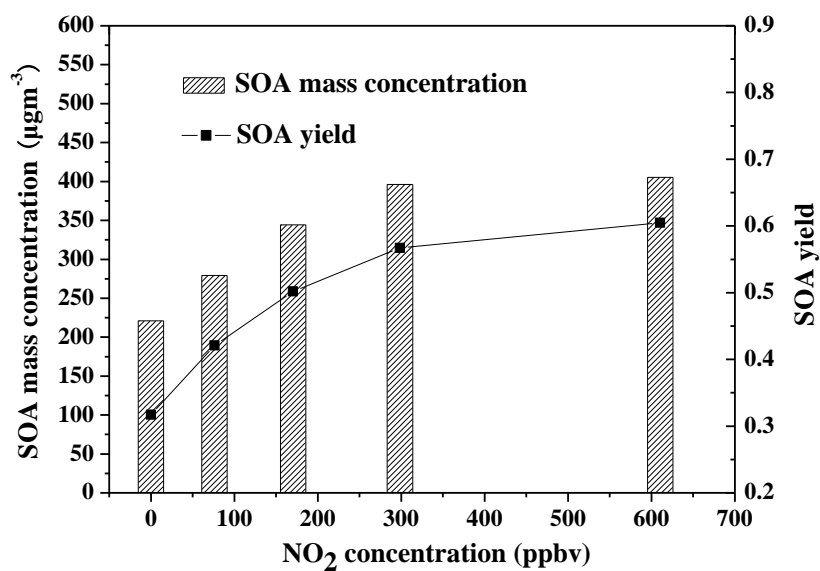
Figure 2. Hu *et al*



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Figure 3. Hu *et al*

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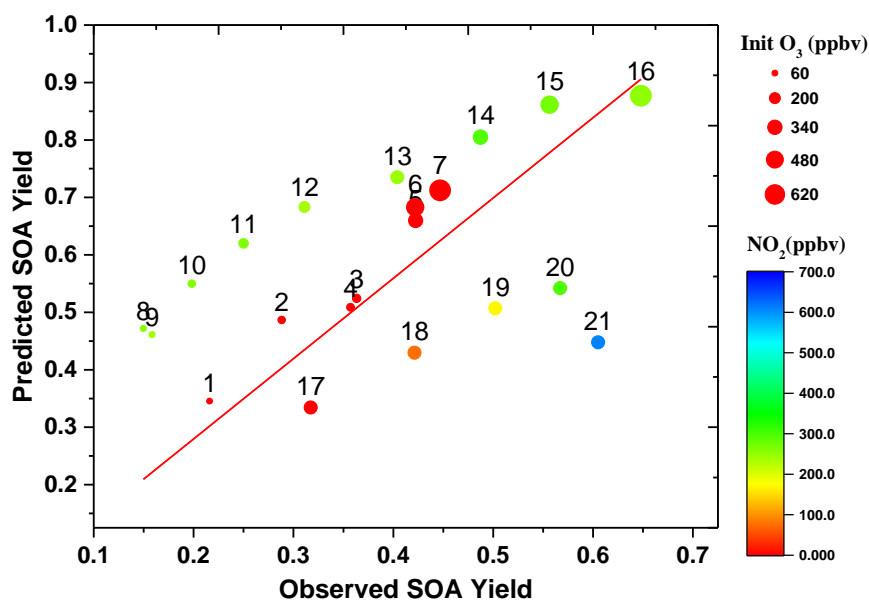
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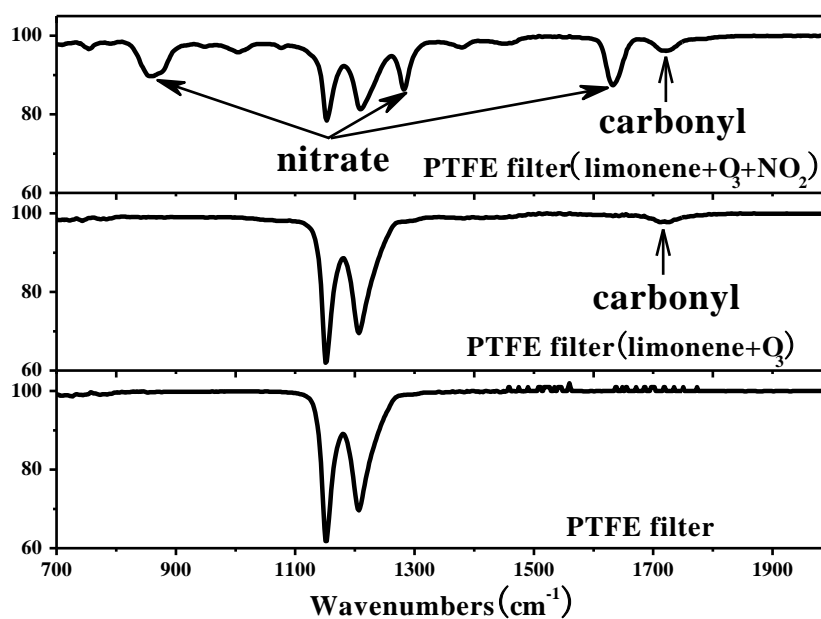
Figure 4. Hu *et al*



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Figure 5. Hu *et al*

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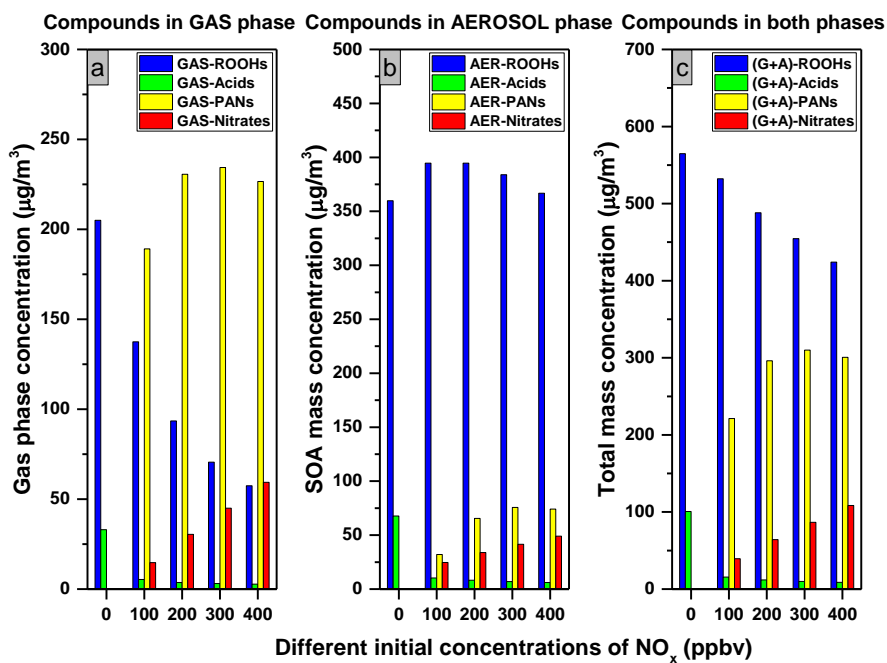
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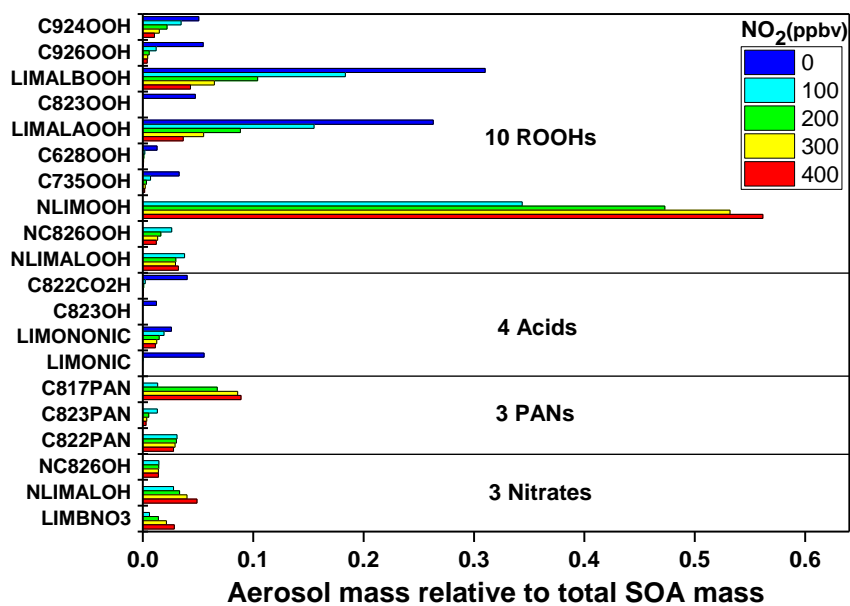
Figure 6. Hu *et al*



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Figure 7. Hu *et al*

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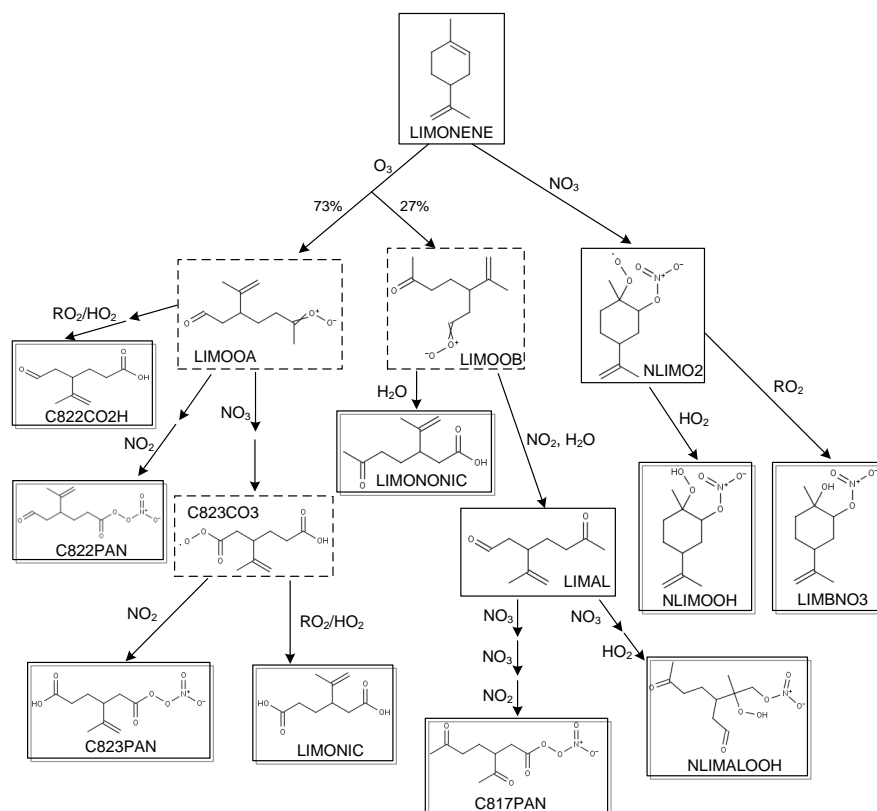
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Scheme 1. Hu *et al*

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1177 **Table 1** Summary of experimental conditions and results in the presence or absence of NO₂^a

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Expt. No.	HC ₀ (ppbv)	[O ₃] ₀ (ppbv)	ΔHC (ppbv)	ΔHC (μg/m ³)	[NO ₂] ₀ (ppbv)	M ₀ (μg/m ³)	SOA Yield
N 1	130	~60	51.1	285	<1 ^c	61.5	0.216
N 2	127	~100	75.7	422	<1 ^c	122	0.288
N 3	123	~120	104	582	<1 ^c	211	0.363
N 4	126	~110	104	581	<1 ^c	207	0.357
N 5	128	~350	128	714	~2 ^c	302	0.422
N 6	121	~500	121	673	~3 ^c	284	0.422
N 7	122	~700	122	680	~3 ^c	304	0.447
N 8	124	~70	53.6	299	256	44.7	0.150
N 9	128	~70	57.1	318	240	50.4	0.158
N 10	128	~100	88.6	494	259	97.9	0.198
N 11	126	~160	116	647	264	162	0.250
N 12	126	~210	126	700	237	218	0.311
N 13	120	~250	120	670	245	271	0.404
N 14	128	~360	128	710	293	346	0.487
N 15	126	~500	126	704	270	392	0.556
N 16	125	~700	125	695	251	451	0.648
N 17 ^b	125	~300	125	697	<2 ^c	221	0.317
N 18 ^b	119	~300	119	663	76.2	279	0.421
N 19 ^b	123	~300	123	686	170	344	0.502
N 20 ^b	125	~300	125	699	299	396	0.567
N 21 ^b	120	~300	120	700	611	405	0.605

1179 ^a Experiments were carried out at temperature of 18.4±1.6°C, and relative humidity of
 1180 12.8% ±2.5%. ^b Experiments were carried out at temperature of 25.4±0.8 °C, and relative humidity
 1181 of 14.5% ±0.7%. ^c the minor NO₂ was produce from corona-discharge when O₃ was produced.

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1191 **Table 2.** The names, structures and formulas of the top 20 condensable compounds

1192 according to their organic aerosol mass fraction.

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Name	Structure	Formula	Name	Structure	Formula
C924OOH		C ₉ H ₁₆ O ₄	C822CO2H		C ₉ H ₁₄ O ₃
C926OOH		C ₉ H ₁₄ O ₅	C823OH		C ₈ H ₁₄ O ₃
LIMALBOOH		C ₁₀ H ₁₆ O ₄	LIMONONIC		C ₁₀ H ₁₆ O ₃
C823OOH		C ₈ H ₁₄ O ₄	LIMONIC		C ₉ H ₁₄ O ₄
LIMALAOOH		C ₁₀ H ₁₆ O ₄	C817PAN		C ₉ H ₁₃ NO ₇
C628OOH		C ₆ H ₁₀ O ₅	C823PAN		C ₉ H ₁₃ NO ₇
C735OOH		C ₇ H ₁₀ O ₅	C822PAN		C ₉ H ₁₃ NO ₆
NLIMOOH		C ₁₀ H ₁₇ NO ₅	NC826OH		C ₈ H ₁₃ NO ₆
NC826OOH		C ₈ H ₁₃ NO ₇	NLIMALOH		C ₁₀ H ₁₇ NO ₆
NLIMALOOH		C ₁₀ H ₁₇ NO ₇	LIMBNO3		C ₁₀ H ₁₇ NO ₄

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