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Influence of temperature on the molecular composition of ions 1

and charged clusters during pure biogenic nucleation 2

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Abstract

47 It was recently shown by the CERN CLOUD experiment that biogenic highly oxygenated 48 molecules (HOMs) form particles under atmospheric conditions in the absence of sulfuric 49 acid where ions enhance the nucleation rate by one to two orders of magnitude. The biogenic HOMs were produced from ozonolysis of α-pinene at 5°C. Here we extend this study to 50 51 compare the molecular composition of positive and negative HOM clusters measured with 52 atmospheric pressure interface time-of-flight mass spectrometers (APi-TOFs), at three 53 different temperatures (25°C, 5°C and -25°C). Most negative HOM clusters include a nitrate 54 (NO₃) ion and the spectra are similar to those seen in the nighttime boreal forest. On the other hand, most positive HOM clusters include an ammonium (NH₄⁺) ion and the spectra are 55 56 characterized by mass bands that differ in their molecular weight by ~20 C atoms, 57 corresponding to HOM dimers. At lower temperatures the average oxygen to carbon (O:C) 58 ratio of the HOM clusters decreases for both polarities, reflecting an overall reduction of 59 HOM formation with decreasing temperature. This indicates a decrease in the rate of 60 autoxidation with temperature due to a rather high activation energy as has previously been 61 determined by quantum chemical calculations. Furthermore, at the lowest temperature 62 (-25°C) the presence of C₃₀ clusters show that HOM monomers start to contribute to the 63 nucleation of positive clusters. These experimental findings are supported by quantum 64 chemical calculations of the binding energies of representative neutral and charged clusters.

1. Introduction

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66 Atmospheric aerosol particles directly affect climate by influencing the transfer of radiant

67 energy through the atmosphere (Boucher et al., 2013). In addition, aerosol particles can

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68 indirectly affect climate, by serving as cloud condensation nuclei (CCN) and ice nuclei (IN).

69 They are of natural or anthropogenic origin, and result from direct emissions (primary

70 particles) or from oxidation of gaseous precursors (secondary particles). Understanding

71 particle formation processes in the atmosphere is important since more than half of the

72 atmospheric aerosol particles may originate from nucleation (Dunne et al., 2016; Merikanto

73 et al., 2009).

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74 Due to its widespread presence and low saturation vapor pressure, sulfuric acid is believed to

75 be the main vapor responsible for new particle formation (NPF) in the atmosphere. Indeed,

76 particle nucleation is dependent on its concentration, albeit with large variability (Kulmala et

al., 2004). The combination of sulfuric acid with ammonia and amines increases nucleation

78 rates due to a higher stability of the initial clusters (Almeida et al., 2013; Kirkby et al., 2011;

79 Kürten et al., 2016). However, these clusters alone cannot explain the particle formation rates

80 observed in the atmosphere. Nucleation rates are greatly enhanced when oxidized organics

81 are present together with sulfuric acid, resulting in NPF rates that closely match those

observed in the atmosphere (Metzger et al., 2010; Riccobono et al., 2014). An important

characteristic of the organic molecules participating in nucleation is their high oxygen content

and consequently low vapor pressure. The formation of these highly oxygenated molecules

85 (HOMs) has been described by Ehn et al. (2014), who found that, following the well-known

86 initial steps of α-pinene ozonolysis through a Criegee intermediate leading to the formation of

87 an RO₂· radical, several repeated cycles of intramolecular hydrogen abstractions and O₂

88 additions produce progressively more oxygenated RO₂· radicals, a mechanism called

89 autoxidation (Crounse et al., 2013). The (extremely) low volatility of the HOMs results in

90 efficient NPF and growth, even in the absence of sulfuric acid (Kirkby et al., 2016; Tröstl et

91 al., 2016). The chemical composition of HOMs during NPF has been identified from α -

92 pinene and pinanediol oxidation by Praplan et al. (2015) and Schobesberger et al. (2013),

93 respectively.

94 Charge has also been shown to enhance nucleation (Kirkby et al., 2011). Ions are produced

95 in the atmosphere mainly by galactic cosmic rays and radon. The primary ions are N^+ , N_2^+ ,

 O^+ , O_2^+ , O_3^+ , O^- and O_2^- (Shuman et al., 2015). These generally form clusters with water

97 (e.g. $(H_2O)H_3O^+$) and after further collisions the positive and negative charges are transferred

98 to trace species with highest and lowest proton affinities, respectively (Ehn et al., 2010). Ions

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are expected to promote NPF by increasing the cluster binding energy and reducing evaporation rates (Hirsikko et al., 2011). Recent laboratory experiments showed that ions increase the nucleation rates of HOMs from the oxidation of α -pinene by one to two orders of magnitude compared to neutral conditions (Kirkby et al. 2016). This is due to two effects, of which the first is more important: 1) an increase in cluster binding energy, which decreases evaporation and 2) an enhanced collision probability, which increases the condensation of polar vapors on the charged clusters (Lehtipalo et al., 2016; Nadykto, 2003).

Temperature plays an important role in nucleation, resulting in strong variations of NPF at different altitudes. Kürten et al. (2016) studied the effect of temperature on nucleation for the sulfuric acid - ammonia system, finding that low temperatures decrease the needed concentration of H_2SO_4 to maintain a certain nucleation rate. Similar results have been found for sulfuric acid – water binary nucleation (Duplissy et al., 2016; Merikanto et al., 2016), where temperatures below $0^{\circ}C$ were needed for NPF to occur at atmospheric concentrations. Up to now, no studies have addressed the temperature effect on NPF driven by HOMs from biogenic precursors such as α -pinene.

In this study we focus on the chemical characterization of the ions and the influence of temperature on their chemical composition during organic nucleation in the absence of sulfuric acid. The importance of such sulfuric acid-free clusters for NPF has been shown in the laboratory (Kirkby et al., 2016; Tröstl et al., 2016) as well as in the field (Bianchi et al., 2016). We present measurements of the NPF process from the detection of primary ions (e.g. N_2^+ , O_2^+ , NO^+) to the formation of clusters in the size range of small particles, all under atmospherically relevant conditions. The experiments were conducted at three different temperatures (-25, 5 and 25 °C) enabling the simulation of pure biogenic NPF representative of different tropospheric altitudes.

2. Methods

2.1. The CLOUD chamber

We conducted experiments at the CERN CLOUD chamber (Cosmics Leaving Outdoor Droplets). With a volume of 26.1 m³, the chamber is built of electropolished stainless steel

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and equipped with a precisely controlled gas system. The temperature inside the chamber is measured with a string of six thermocouples (TC, type K) which were mounted horizontally between the chamber wall and the center of the chamber at distances of 100, 170, 270, 400, 650, and 950 mm from the chamber wall (Hoyle et al., 2016). The temperature is controlled accurately (with a precision of \pm 0.1°C) at any tropospheric temperature between -65 and 30 °C (in addition, the temperature can be raised to 100 °C for cleaning). The chamber enables atmospheric simulations under highly stable experimental conditions with low particle wall loss and low contamination levels (more details of the CLOUD chamber can be found in Kirkby et al. (2011) and Duplissy et al. (2016)). Before the start of the experiments the CLOUD chamber was cleaned by rinsing the walls with ultra-pure water, followed by heating to 100°C and flushing at a high rate with humidified synthetic air and elevated ozone (several ppmv) (Kirkby et al., 2016). This resulted in SO₂ and H₂SO₄ concentrations that were below the detection limit (<15 pptv and <5 × 10^4 cm⁻³, respectively), and total organics (largely comprising high volatility C₁–C₃ compounds) that were below 150 pptv.

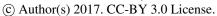
The air in the chamber is ionized by galactic cosmic rays (GCR); higher ion generation rates can be induced by a pion beam (π^+) from the CERN Proton Synchrotron enabling controlled simulation of galactic cosmic rays throughout the troposphere. Therefore, the total ion-pair production rate in the chamber is between 2 (no beam) and 100 cm⁻³ s⁻¹ (maximum available beam intensity, Franchin et al., 2015).

2.2. Instrumentation

The main instruments employed for this study were atmospheric pressure interface time-of-flight (APi-TOF, Aerodyne Research Inc. & Tofwerk AG) mass spectrometers. The APi-TOF is able to measure the intensity of positive or negative ions and cluster ions over a wide range of mass-to-charge ratios at concentrations relevant for the ambient atmosphere. The instrument has two main parts. The first is the atmospheric pressure interface (APi) where ions are transferred from atmospheric pressure to low pressures via three differentially pumped vacuum stages. Ions are focused and guided by two quadrupoles and ion lenses. The second is the time-of-flight mass analyzer (TOF), where the pressure is approximately 10^{-6} mbar. The sample flow from the chamber was 10 L/min and the core-sampled flow into the APi was 0.8 L/min, with the remaining flow being discarded.

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We calibrated the APi-TOF using trioctylmethylammonium bis (trifluoromethylsulfonyl) imide (TBMA, $C_{27}H_{54}F_6N_2O_4S_2$) to facilitate the exact ion mass determination in both positive and negative ion modes. We employed two calibration methods, the first one by nebulizing TBMA and separating cluster ions with a high-resolution ultra-fine differential mobility analyzer (UDMA) (see Steiner et al. (2014) for more information); the second one by using electrospray ionization of a TBMA solution. The calibration with the electrospray ionization was performed three times, one for each temperature. These calibrations enabled mass/charge (m/z) measurements with high accuracy up to 1500 Th in the positive ion mode and 900 Th in the negative ion mode.

Additionally, two peaks in the positive ion mode were identified as contaminants and also used for calibration purposes at the three different temperatures: $C_{10}H_{14}OH^{+}$ and $C_{20}H_{28}O_{2}H^{+}$. These peaks were present before the addition of ozone in the chamber (therefore being most likely not products of α -pinene ozonolysis) and were also detected by a proton transfer reaction time of flight mass spectrometer (PTR-TOF-MS). Both peaks appeared at the same m/z at all three temperatures. Therefore, based on the calibrations with the UDMA, the electrospray and the two organic calibration peaks, we expect an accurate mass calibration at the three temperatures.

2.3. Experimental conditions

All ambient ion composition data reported here were obtained during nucleation experiments from pure α -pinene ozonolysis. The experiments were conducted under dark conditions, at a relative humidity (RH) of 38% with an O_3 mixing ratio between 33 and 43 ppbv (Table 1). The APi-TOF measurements were made under both galactic cosmic ray (GCR) and π^+ beam conditions, with ion-pair concentrations around 700 cm⁻³ and 4000 cm⁻³, respectively.

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Table 1. Experiments performed at the CLOUD chamber.

Campaign	Experiment No.	Ionization	α-pinene (pptv)	O_3 (ppbv)	Mass spectrometer polarity	Temperature (°C)
CLOUD 8	1211.02	GCR	258	33.8	Negative	5
CLOUD 10	1710.04	π^{+} beam	618	41.5	Positive	5
CLOUD 10	1712.04	π^{+} beam	511	40.3	Negative and positive	25
CLOUD 10	1727.04	π^{+} beam	312	43.3	Negative and positive	-25

2.4. Quantum chemical calculations

Quantum chemical calculations were performed on the cluster ion formation from the oxidation products of α -pinene. The Gibbs free energies of formation of representative HOM clusters were calculated using the MO62X functional (Zhao and Truhlar, 2008), and the 6-31+G(d) basis set (Ditchfield, 1971) using the Gaussian09 program (Frisch et al., 2009). This method has been previously applied for clusters containing large organic molecules (Kirkby et al., 2016).

3. Results and discussion

3.1. Ion composition

Under relatively dry conditions, the main detected positive ions were N_2H^+ and O_2^+ . With increasing RH we observed the water clusters H_3O^+ , $(H_2O)\cdot H_3O^+$ and $(H_2O)_2\cdot H_3O^+$ as well as NH_4^+ , $C_5H_5NH^+$ (protonated pyridine), Na^+ , and K^+ . The concentrations of the precursors of some of the latter ions are expected to be very low: for example, NH_3 mixing ratios were previously found to be in the range of 0.3 pptv (at -25 °C), 2 pptv (at 5 °C) and 4.3 pptv (at 25 °C) (Kürten et al., 2016). For the negative ions, NO_3^- was the main detected background signal. Before adding any trace gas to the chamber the signal of HSO_4^- was at a level of 1% of the NO_3^- signal (corresponding to $<5\cdot10^{-4}$ molecules cm⁻³, Kirkby et al., 2016), excluding any contribution of sulfuric acid to nucleation in our experiments.

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After initiating α -pinene ozonolysis, more than 460 organic ions were identified in the positive spectrum. The majority of peaks were clustered with NH₄⁺, while only 10.2 % of the identified peaks were composed of protonated organic molecules. In both cases the organic core was of the type $C_{7-10}H_{10-16}O_{1-10}$ for the monomer region and $C_{17-20}H_{24-32}O_{5-19}$ for the dimer region.

In the negative spectrum we identified more than 530 HOMs, of which \sim 62% corresponded to organic clusters with NO₃ or, to a lesser degree, HNO₃· NO₃. The rest of the peaks were negatively charged organic molecules. In general, the organic core of the molecules was of the type $C_{7-10}H_{9-16}O_{3-12}$ in the monomer region and $C_{17-20}H_{19-32}O_{10-20}$ in the dimer region. For brevity we refer to the monomer, dimer (and n-mer) as C_{10} , C_{20} and $C_{(10n)}$ respectively. Here, the subscript indicates the maximum number of carbon atoms in these molecules, even though the bands include species with slightly fewer carbon atoms.

3.1.1. Positive spectrum

The positive spectrum is characterized by bands of high intensity at C_{20} intervals, as shown in Figure 1B. Although we detected the monomer band (C_{10}), its integrated intensity was much lower than the C_{20} band; furthermore, the trimer and pentamer bands were almost completely absent. Based on chemical ionization mass spectrometry measurements, Kirkby et al. (2016) calculated that the HOM molar yield at 5°C was 3.2% for the ozonolysis of α -pinene, with a fractional yield of 10 to 20% for dimers. A combination reaction of two oxidized peroxy radicals has been previously reported to explain the rapid formation of dimers resulting in covalently bound molecules (see Section 3.3). The pronounced dimer signal with NH_4^+ indicates that (low-volatility) dimers are necessary for positive ion nucleation and initial growth. We observe growth by dimer steps up to C_{80} and possibly even C_{100} . A cluster of two dimers, C_{40} , with a mass/charge in the range of ~ 700 - 1100 Th, has a mobility diameter around 1.5 nm (based on Ehn et al. (2011)).

Our observation of $HOMs-NH_4^+$ clusters implies strong hydrogen bonding between the two species. This is confirmed by quantum chemical calculations which shall be discussed in Section 3.3. Although hydrogen bonding could also be expected between HOMs and H_3O^+ , we do not observe such clusters. This probably arises from the higher proton

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affinity of NH₃, 203.6 kcal/mol, compared with H₂O, 164.8 kcal/mol (Hunter and Lias, 1998). Thus, most H₃O⁺ ions in CLOUD will transfer their proton to NH₃ to form NH₄⁺.

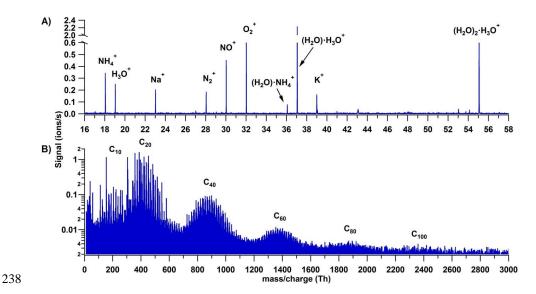


Figure 1. Positive spectra at 5° C. A) Low mass region, where primary ions from galactic cosmic ray are observed, as well as secondary ions such as NH_4^+ which are formed by charge transfer to contaminants. B) Higher mass region during pure biogenic nucleation, which shows broad bands in steps of C_{20} . Most of the peaks represent clusters with NH_4^+ .

3.1.2. Negative spectrum

In the negative spectra, the monomer, dimer and trimer bands are observed during nucleation (Fig. 2). Monomers and dimers have similar signal intensities, whereas the trimer intensity is at least 10 times lower (Figure 2A and B). The trimer signal is reduced since it is a cluster of two gas phase species ($C_{10}+C_{20}$). Additionally, a lower transmission in the APi-TOF may also be a reason for the reduced signal.

In Fig. 2, we compare the CLOUD negative-ion spectrum with the one from nocturnal atmospheric measurements from the boreal forest at Hyytiälä as reported by Ehn et al. (2010). Panels 2A and 2B show the negative spectrum of α -pinene ozonolysis in the CLOUD chamber on logarithmic and linear scales, respectively. Panel 2C shows the Hyytiälä spectrum for comparison. Although the figure shows unit mass resolution, the high resolution

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analysis confirms the identical composition for the main peaks: $C_8H_{12}O_7 \cdot NO_3^-$, $C_{10}H_{14}O_7 \cdot NO_3^-$, $C_{10}H_{14}O_9 \cdot NO_3^-$, $C_{10}H_{16}O_{10} \cdot NO_3^-$ and $C_{10}H_{14}O_{11} \cdot NO_3^-$ (marked in the monomer region), and $C_{19}H_{28}O_{11} \cdot NO_3$, $C_{20}H_{32}O_{13} \cdot NO_3^-$, $C_{19}H_{28}O_{13} \cdot NO_3^-$, $C_{20}H_{30}O_{14} \cdot NO_3^-$, $C_{20}H_{30}O_{16} \cdot NO_3^-$ and $C_{20}H_{30}O_{18} \cdot NO_3^-$ (marked in the dimer region). The close correspondence in terms of composition of the main HOMs from the lab and the field both in the monomer and dimer region indicates a close reproduction of the atmospheric night-time conditions at Hyytiälä by the CLOUD experiment. In both cases the ion composition was dominated by HOMs clustered with NO_3^- . However, Ehn et al. (2010) did not report nocturnal nucleation, possibly because of a higher ambient condensation sink than in the CLOUD chamber.

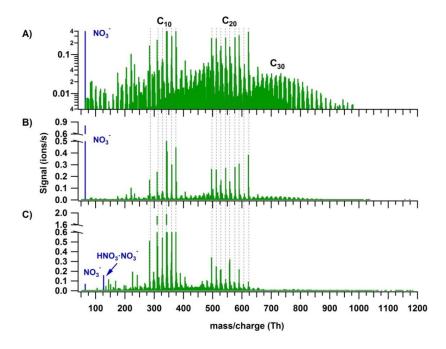


Figure 2. Comparison of the negative ion composition during α-pinene ozonolysis in CLOUD and during night time in the boreal forest at Hyytiälä (Finland). A) CLOUD spectrum on a logarithmic scale. B) CLOUD spectrum on a linear scale. C) Typical night time spectrum from the boreal forest at Hyytiälä (Finland), adapted from Ehn et al. (2010).

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3.2. Temperature dependence

Experiments at three different temperatures (25 °C, 5°C and -25 °C) were conducted at similar relative humidity and ozone mixing ratios (Table 1 and Figure 3). Mass defect plots are shown for the same data in Figure 4. The mass defect is the difference between the exact and the integer mass and is shown on the y-axis versus the mass/charge on the x-axis. Each point represents a distinct atomic composition of a molecule or cluster. Although the observations described in the following are valid for both polarities, the positive ion mode shows the differences in the chemistry at the three temperatures more clearly.

The first point to note is the change in the distribution of the signal intensity seen in Figure 3 (height of the peaks) and in Figure 4 (size of the dots) with temperature. In the positive ion mode, the dimer band has the highest intensity at 25 and 5°C, while at -25°C the intensity of the monomer becomes comparable to that of the dimer. This indicates a reduced rate of dimer formation at -25 °C, or that the intensity of the ion signal depends on both the concentration of the neutral compound and on the stability of the ion cluster. Although the monomer concentration is higher than that of the dimers (Tröstl et al., 2016), the C_{20} ions are the more stable ion clusters as they can form more easily two hydrogen bonds with NH_4^+ (see Section 3.3). Thus, positive clusters formed from monomers may not be stable enough at higher temperatures. Moreover, charge transfer to dimers is also favored.

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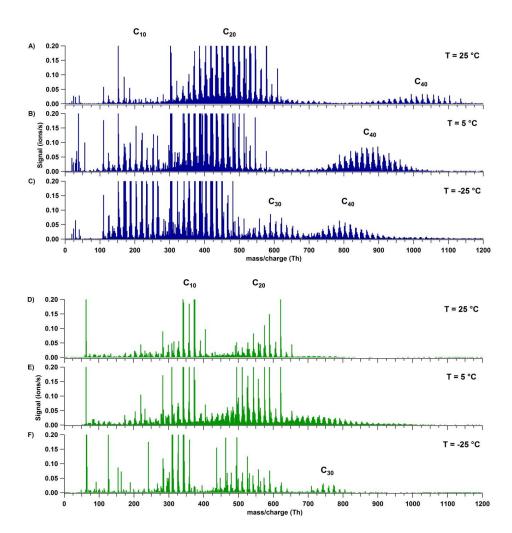


Figure 3. Positive (A-C) and negative (D-F) mass spectra during pure biogenic nucleation induced by ozonolysis of \Box -pinene) at three temperatures: 25°C (A, D), 5°C (B, E) and -25°C (C, F). A progressive shift towards a lower oxygen content and lower masses is observed in all bands as the temperature decreases. Moreover, the appearance of C_{30} species can be seen in the positive spectrum at the lowest temperature (C).

The data also show a "shift" in all band distributions towards higher masses with increasing temperature, denoting a higher concentration of the more highly oxygenated molecules and the appearance of progressively more oxygenated compounds at higher temperatures. The shift is even more pronounced in the higher mass bands, as clearly seen in

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the C_{40} band of the positive ion mode in Figure 3 (A-C). In this case the combination of two HOM dimers to a C_{40} cluster essentially doubles the shift of the band towards higher mass/charge at higher temperatures compared to the C_{20} band. Moreover, the width of each band increases with temperature, as clearly seen in the positive ion mode in Figure 4, especially for the C_{40} band. At high temperatures, the production of more highly oxygenated HOMs seems to increase the possible combinations of clusters, resulting in a wider band distribution.

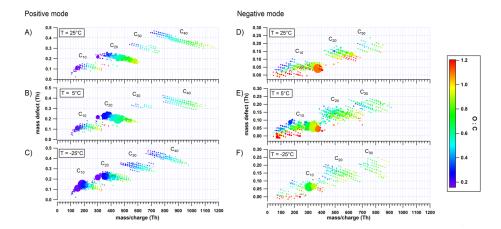


Figure 4. Mass defect plots with the color code denoting the O:C ratio (of the organic core) at 25, 5 and -25°C for positive (A-C) and negative ion mode (D-F). A lower O:C ratio is observed in the positive ion mode than in the negative ion mode. The intensity of the main peaks (linearly proportional to the size of the dots) changes with temperature for both polarities due to a lower degree of oxygenation at lower temperature.

This trend in the spectra indicates that the unimolecular autoxidation reaction accelerates at higher temperatures in competition to the bimolecular termination reactions with HO₂ and RO₂. This is expected. If unimolecular and bimolecular reactions are competitive, the unimolecular process will have a much higher barrier because the pre-exponential term for a unimolecular process is a vibrational frequency while the pre-exponential term for the bimolecular process is at most the bimolecular collision frequency, which is four orders of magnitude lower. Quantum chemical calculations determine

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activation energies between 22.56 and 29.46 kcal/mol for the autoxidation of different RO_2 radicals from α -pinene (Rissanen et al., 2015). Thus, such a high barrier will strongly reduce the autoxidation rate at the low temperatures.

The change in the rate of autoxidation is also reflected in the O:C ratio, both in the positive ion mode (Figure 4 A-C), and the negative ion mode (D-F), showing a clear increase with increasing temperature. The average O:C ratios (weighted by the peak intensities) are presented in Table 2 for both polarities and the three temperatures, for all the identified peaks (total) and separately for the monomer and dimer bands For a temperature change from 25 to -25° C the O:C ratio decreases for monomers, dimers and total number of peaks. At high masses (e.g., for the C₃₀ and C₄₀ bands), the O:C ratio may be slightly biased since accurate identification of the molecules is less straightforward: as an example, C₃₉H₅₆O₂₅·NH₄⁺ has an exact mass of 942.34 Th (O/C = 0.64), which is very similar to C₄₀H₆₀O₂₄·NH₄⁺ at 942.38 Th (O/C = 0.60). However, such possible misidentification would not influence the calculated total O/C by more than 0.05, and the main conclusions presented here remain robust.

Table 2. Signal weighted average O:C ratios for positive and negative spectra at 25, 5 and -25 °C.

	O/C						
Temperature (°C)	Positive mode			Negative mode			
	Monomer	Dimer	Total	Monomer	Dimer	Total	
25	0.37	0.57	0.54	0.94	0.81	0.90	
5	0.34	0.51	0.49	0.88	0.66	0.75	
-25	0.31	0.38	0.36	0.79	0.65	0.68	

The O:C ratios are higher for the negative ions than for the positive ions at any of the three temperatures. Although some of the organic cores are the same in the positive and negative ion mode, the intensity of the peaks of the most oxygenated species is higher in the negative spectra. While the measured O:C ratio ranges between 0.4 and 1.2 in the negative ion mode, it is between 0.1 and 1.2 in the positive ion mode. An O:C ratio of 0.1, which was detected only in the positive ion mode, corresponds to monomers with 1 oxygen atom or dimers with two oxygen atoms. The presence of molecules with such low oxygen content was

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also confirmed with a proton transfer reaction time-of-flight mass spectrometer (PTR-TOF-MS), at least in the monomer region. The ions with this low O:C ratio are probably from the known main oxidation products like pinonaldehyde, pinonic acid, etc. It is likely that these molecules, which were detected only in the positive mode, contribute only to the growth of the newly formed particles (if at all) rather than to nucleation, owing to their high volatility (Tröstl et al., 2016). In this sense, the positive spectrum could reveal both the molecules that participate in the new particle formation and those that contribute to growth. The differences in the O:C ratios between the two polarities are a result of the affinities of the organic molecules to form clusters either with NO₃ or NH₄, which, in turn, depends on the molecular structure and the functional groups. Hyttinen et al. (2015) reported the binding energies of selected highly oxygenated products of cyclohexene detected by a nitrate CIMS, finding that the addition of OOH groups to the HOM strengthens the binding of the organic core with NO₃. Even when the number of H-bonds between NO₃ and HOM remains the same, the addition of more oxygen atoms to the organic compound could strengthen the bonding with the NO₃ ion. Thus, the less oxygenated HOMs were not detected in those experiments, neither in ours, in the negative mode. The binding energies were calculated for the positive mode HOMs-NH₄⁺ and are discussed in Section 3.3.

We also tested to which extent the formation of the C_{40} band could be reproduced by permutation of the potential C_{20} clusters weighted by the dimer signal intensity. Figure 5 shows the measured spectrum (blue) and two types of modeled tetramers: one combining all peaks from the C_{20} band (light gray) and one combining only those peaks with an organic core with $O/C \ge 0.4$, i.e. likely non-volatile molecules (dark gray). The better consistency of the latter with the measured tetramer band suggests that only the molecules with $O/C \ge 0.4$ are able to form the tetramer cluster. This would mean that C_{20} molecules with 2-7 oxygen atoms are likely not to contribute to the nucleation, but only to the growth of the newly formed particles.

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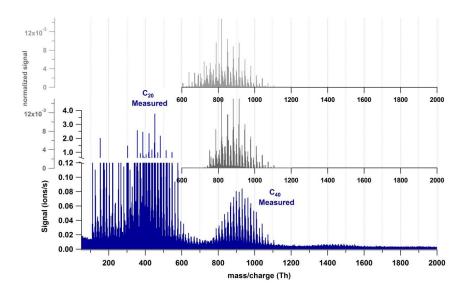


Figure 5. Comparison of the positive ion mode spectrum measured (blue), the C_{40} band obtained by the combination of all C_{20} molecules (light gray) and the C_{40} band obtained by combination of only the C_{20} molecules with $O/C \ge 0.4$ (dark gray). The low or absent signals at the lower masses obtained by permutation suggests that only the highly oxygenated dimers are able to cluster and form C_{40} .

These two observations (change in signal distribution and band "shift") are not only valid for positive and negative ions, but also for the neutral molecules as observed by two nitrate chemical-ionization atmospheric-pressure-interface time-of-flight mass spectrometers (CI-APi-TOF; Aerodyne Research Inc. and Tofwerk AG). This confirms that there is indeed a change in the HOM composition with different temperature rather than a charge redistribution effect which would only be observed for the ions (APi-TOF). The detailed analysis of the neutral molecules detected by these CI-APi-TOFs will be subject of another paper and is not discussed here.

A third distinctive trend in the positive mode spectra at the three temperatures is the increase in signal intensity of the C₃₀ band at -25°C. The increase in the signal of the trimer also seems to occur in the negative ion mode when comparing panels D and F in Figure 3. For this polarity, data from two campaigns were combined (Table 1). To avoid a bias by possible differences in the APi-TOF settings, we only compare the temperatures from the same campaign, CLOUD 10, therefore experiments at 25 °C and -25 °C. The increase in the

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trimer signal may be due to greater stability of the monomer-dimer clusters or even of three C₁₀ molecules at low temperatures, as further discussed below.

3.3. Quantum chemical calculations

Three points were addressed in the quantum chemical calculations to elucidate the most likely formation pathway for the first clusters, and its temperature dependence. These included (i) the stability of the organic cores with NO_3^- and NH_4^+ depending on the binding functional group, (ii) the difference between charged and neutral clusters in terms of clustering energies, and finally (iii) the possible nature of clusters in the dimer and trimer region.

The calculations showed that among the different functional groups the best interacting groups with NO_3^- are in order of importance carboxylic acids (R–C(=O)–OH), hydroxyls (R–OH), peroxy acids (R–C(=O)–O–OH), hydroperoxides (R–O–OH) and carbonyls (R–(R'–) C=O). On the other hand, NH_4^+ preferably forms a hydrogen bond with the carbonyl group independent of which functional group the carbonyl group is linked to: Figure 6 shows examples of NH_4^+ clusters with corresponding free energies of formation for carbonyls (ΔG = -17.98 kcal/mol), carboxylic acid (ΔG = -17.32 kcal/mol), and peroxy acid (ΔG = -17.46 kcal/mol). For the three examples shown, the interaction of one hydrogen from NH_4^+ with a C=O group is already very stable with a free energy of cluster ion formation close to -18 kcal/mol.

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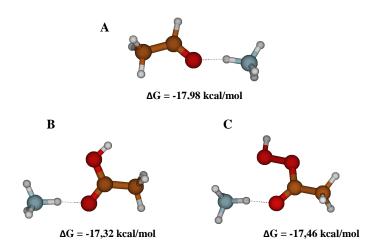


Figure 6. Quantum chemical calculations of the free energy related to the cluster formation between $\mathrm{NH_4}^+$ and three structurally similar molecules with different functional groups: A) acetaldehyde, B) acetic acid and C) peracetic acid.

To evaluate the effect of the presence of a second C=O to the binding of the organic compound with NH₄⁺, we performed a series of calculations with a set of surrogates containing two C=O groups separated by a different number of atoms, as shown in Figure 7. The addition of a second functional group allows the formation of an additional hydrogenbond, increasing the stability of the cluster considerably (almost two folds) from about -18 kcal/mol to -34.07 kcal/mol, whereby the position of the second functional group to form an optimal hydrogen bond (with a 180° angle for N-H-O) strongly influences the stability of the cluster, as can be seen in Figure 7. Thus, optimal separation and conformational flexibility of functional groups is needed to enable an effective formation of two hydrogen bonds with NH₄⁺. This could be an explanation for the observation that the signal intensity is higher for dimers than for monomers, as dimers can more easily form two optimal hydrogen bonds with NH₄⁺.

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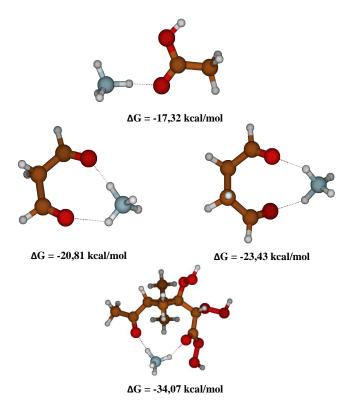


Figure 7. Quantum chemical calculations for different organic molecules with a carbonyl as the interacting functional group with NH_4^+ . Increasing the interacting groups from one to two increases the stability of the cluster. The distance between the interacting groups also influences the cluster stability.

As shown by Kirkby et al. (2016), ions increase the nucleation rates by one to two orders of magnitudes compared to neutral nucleation. This is expected due to the strong electrostatic interaction between charged clusters. To understand how the stability difference relates to the increase in the nucleation rate, the ΔGs of charged and neutral clusters were compared. For this, $C_{10}H_{14}O_7$ and $C_{20}H_{30}O_{14}$ were selected as representative molecules of the monomer and dimer region, respectively (Kirkby et al., 2016). Table 3 shows the calculated free energies of formation (ΔG) of neutral, positive and negative clusters from these C_{10} and C_{20} molecules at the three temperatures of the experiment. Results show that at 5°C, for

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example, ΔG of the neutral dimer ($C_{10} + C_{10}$) is -5.76 kcal/mol while it decreases to -20.95 kcal/mol when a neutral and a negative ion form a cluster ($C_{10} + C_{10}$). Similarly, trimers show a substantial increase in stability when they are charged, i.e., from -2.15 kcal/mol to -19.9 kcal/mol, for the neutral and negative cases, respectively. The reduced values of ΔG for the charged clusters (positive and negative) indicate a substantial decrease in the evaporation rate compared to that for neutral clusters, and, therefore, higher stability. Comparing the NH₄⁺ and NO₃⁻ clusters, the energies of formation for the monomer are -22.5 kcal/mol and -25.99 kcal/mol, respectively, showing slightly higher stability for the negative cluster. Inversely, the covalently bound dimer showed greater stability for the positive ion (-30.9 kcal/mol) compared to the negative ion (-25.65 kcal/mol).

Table 3. Gibbs free energies of cluster formation ΔG at three different temperatures. ΔG for the molecules $C_{10}H_{14}O_7$ (C_{10}) and $C_{20}H_{30}O_{14}$ (C_{20}) forming neutral, as well as negative and positive ion clusters.

	Cluster process	$\Delta G_{-25^{\circ}C}$ (kcal/mol)	$\Delta G_{5^{\circ}C}$ (kcal/mol)	$\Delta G_{25^{\circ}C}$ (kcal/mol)
Neutral	$C_{10} + C_{10}$	-7.33	-5.76	-4.70
	$C_{10} + C_{20}$	-3.28	-2.15	-1.39
Positive	$C_{10} + N{H_4}^+$	-23.40	-22.50	-21.80
	$C_{20} + NH_4^{+}$	-31.80	-30.90	-30.20
	$C_{10} + C_{10} \cdot NH_4^{\ +}$	-12.90	-11.70	-10.90
	$C_{20} + C_{10} \cdot NH_4^{\ +}$	-26.00	-24.30	-23.30
	$C_{10} + C_{20} \cdot NH_4^{+}$	-17.60	-15.90	-14.80
Negative	$C_{10} + C_{10}$	-22.22	-20.95	-20.09
	$C_{20} + C_{10}$	-21.36	-19.90	-18.91
	$C_{10} + NO_3$	-27.27	-25.99	-25.14
	$C_{20} + NO_3$	-26.97	-25.65	-24.75
	$C_{10} + C_{10} \cdot NO_3$	-11.34	-10.09	-9.25

 The temperature dependence of cluster formation is shown in Figure 8 for the positive ion clusters. The blue and brown solid lines represent the needed ΔG for evaporation-collision equilibrium at 0.3 pptv and 1 pptv HOM mixing ratio, respectively, calculated as described by Ortega et al. (2012). The markers show the calculated formation enthalpies ΔG for each of the possible clusters. For all cases, the trend shows an evident decrease in ΔG with decreasing temperature, with a correspondingly reduced evaporation rate.

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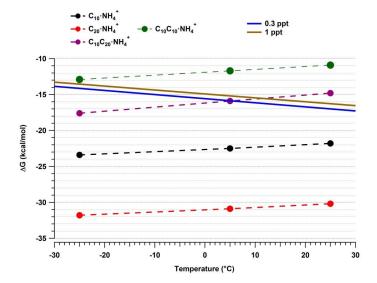


Figure 8. Quantum chemical calculations of Gibbs free energies for cluster formation at -25, 5 and 25°C. Solid lines represent the required ΔG for equilibrium between evaporation and collision rates at 0.3 pptv and 1 pptv of the HOM mixing ratio, respectively. Markers show the ΔG for each cluster (organic core clustered with NH_4^+) at the three temperatures. $C_{10} \cdot NH_4^+$ (black circles) represent the monomer, $C_{20} \cdot NH_4^+$ (red circles) represent the covalently bound dimer, $C_{10}C_{10} \cdot NH_4^+$ (green circles) represent the dimer formed by the clustering of two monomers and $C_{10}C_{20} \cdot NH_4^+$ (purple circles) denote the preferential pathway for the trimer cluster (see Table 3).

At all three temperatures, the monomer cluster $C_{10} \cdot NH_4^+$ falls well below the equilibrium lines, indicating high stability. Even though the difference between -25°C and 25°C is just -1.6 kcal/mol in free energy, it is enough to produce a substantial difference in the intensity of the band, increasing the signal at least 8-fold at -25°C (as discussed in Section 3.2). In the case of the dimers, we consider the possibility of their formation by collision of a monomer $C_{10} \cdot NH_4^+$ with another C_{10} (resulting in a $C_{10}C_{10} \cdot NH_4^+$ cluster) or the dimer as $C_{20} \cdot NH_4^+$ cluster. The calculations show clearly that the cluster $C_{10}C_{10} \cdot NH_4^+$ is not stable at any of the three temperatures (green line). In contrast, the covalently-bound C_{20} forms very stable positive and negative ion clusters (see Table 3). Trimers are mainly observed at lower temperatures. Since the $C_{10}C_{10} \cdot NH_4^+$ cluster is not very stable, we discard the possibility of a trimer formation of the type $C_{10}C_{10}C_{10} \cdot NH_4^+$. Thus, the trimer is likely the combination of a monomer and a covalently-bound dimer $(C_{20}C_{10} \cdot NH_4^+)$. According to our calculations (Table

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3) the preferred evaporation path for this cluster is the loss of C_{10} rather than the evaporation of C_{20} . Therefore, we have chosen to represent only this path in Figure 8. The \Box G of this cluster crosses the evaporation-condensation equilibrium around 5 °C and 14 °C for a HOM mixing ratio of 0.3 pptv and 1 pptv, respectively, in good agreement with the observed signal increase of the trimer at -25 °C (Figure 3 A-C). It is important to note that, due to the uncertainty in the calculations, estimated to be \leq 2 kcal/mol, we do not consider the crossing as an exact reference.

The ΔG of the negative ion clusters, which are also presented in Table 3, decrease similarly to the positive ion clusters by around 2 kcal/mol between 25°C and -25°C. The cluster formation energies of the monomer and the dimer with NO_3^- are in agreement with the observed comparable signal intensity in the spectrum (Figure 2) in a similar way as the positive ion clusters. The covalently-bonded dimer ion $C_{20} \cdot NO_3^-$ is also more stable compared to the dimer cluster $C_{10}C_{10} \cdot NO_3^-$, suggesting that the observed composition results from covalently bonded dimers clustering with NO_3^- rather than two individual C_{10} clustering to form a dimer.

The formation of a covalently bonded trimer seems unlikely, so the formation of highly oxygenated molecules is restricted to the monomer and dimer region. The trimer could result from the clustering of C_{10} and C_{20} species. Similarly, and based on the C_{20} pattern observed in Figure 1B, we believe that the formation of the tetramer corresponds to the collision of two dimers. No calculations were done for this case due to the complexity related to the sizes of the molecules, which prevents feasible high level quantum chemical calculations.

Finally, the comparison between the ΔG values for charged and neutral clusters as presented in Table 3 confirms the expected higher stability of ion clusters, decreasing the evaporation rate of the nucleating clusters and enhancing new particle formation.

4. Conclusions

Ions observed during pure biogenic ion-induced nucleation were comprised of mainly organics clustered with NO₃⁻ and NH₄⁺ and to a lesser extent charged organic molecules only or organics clustered with HNO₃NO₃⁻. We found good correspondence between the negative

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ions measured in CLOUD with those observed in the boreal forest of Hyytiälä. The observed similarity in the composition of the HOMs in the monomer and dimer region during newparticle formation experiments at CLOUD suggests that pure biogenic nucleation might be possible during night time if the condensation sink is sufficiently low, i.e., comparable to that in the CLOUD chamber, where the wall loss rate for H₂SO₄ is 1.8·10⁻³ s⁻¹ (Kirkby et al., 2016). The positive mass spectrum showed a distinctive pattern corresponding to progressive addition of dimers (C₂₀), up to cluster sizes in the range of stable small particles.

Temperature strongly influenced the composition of the detected molecules in several ways. With increasing temperature, a higher oxygen content (O:C ratio) in the molecules was observed in both the positive and the negative mode. This indicates an increase in the autoxidation rate of peroxy radicals which is in competition with their bimolecular termination reactions with HO₂ and RO₂.

A broader range of organic molecules was found to form clusters with NH_4^+ than with NO_3^- . Quantum chemical calculations using simplified molecules show that NH_4^+ preferably forms a hydrogen bond with a carbonyl group independently of other functional groups nearby. The addition of a second hydrogen bond was found to increase the cluster stability substantially. Thus, the C_{20} -ions are the more stable ion clusters as they can form more easily two hydrogen bonds with NH_4^+ . Although molecules with low oxygen content were measured in the C_{20} band (1 - 4 oxygen atoms), only the molecules with $O/C \ge 0.4$ seem to be able to combine to form larger clusters.

The quantum chemical calculations showed that the covalently-bonded dimer C_{20} · NO_3^- is also more stable than the dimer cluster $C_{10}C_{10}$ · NO_3^- , suggesting that the observed composition results from covalently bonded molecules clustering with NO_3^- rather than C_{10} clusters.

Temperature affected cluster formation by decreasing evaporation rates at lower temperatures, despite of the lower O:C ratio. In the positive mode a pronounced growth of clusters by addition of C_{20} -HOMs was observed. The formation of a C_{30} -cluster only appeared at the lowest temperature, which was supported by quantum chemical calculations. In the negative mode it appeared as well that the signal of the C_{30} -clusters became stronger with lower temperature. The C_{40} - and higher clusters were probably not seen because of too

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549 low sensitivity in this mass range due to the applied instrumental settings. More

550 measurements are needed to determine if the cluster growth of positive and negative ions

551 proceeds in a similar or different way.

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578 **6. References**

- 579 Almeida, J., Schobesberger, S., Kürten, A., Ortega, I. K., Kupiainen-Määttä, O., Praplan, A.
- 580 P., Adamov, A., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J.,
- 581 Donahue, N. M., Downard, A., Dunne, E., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin,
- A., Guida, R., Hakala, J., Hansel, A., Heinritzi, M., Henschel, H., Jokinen, T., Junninen, H.,
- 583 Kajos, M., Kangasluoma, J., Keskinen, H., Kupc, A., Kurtén, T., Kvashin, A. N., Laaksonen,
- 584 A., Lehtipalo, K., Leiminger, M., Leppä, J., Loukonen, V., Makhmutov, V., Mathot, S.,
- McGrath, M. J., Nieminen, T., Olenius, T., Onnela, A., Petäjä, T., Riccobono, F., Riipinen, I.,
- 586 Rissanen, M., Rondo, L., Ruuskanen, T., Santos, F. D., Sarnela, N., Schallhart, S.,
- 587 Schnitzhofer, R., Seinfeld, J. H., Simon, M., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé,
- 588 A., Tröstl, J., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y., Virtanen, A., Vrtala, A.,
- Wagner, P. E., Weingartner, E., Wex, H., Williamson, C., Wimmer, D., Ye, P., Yli-Juuti, T.,
- 590 Carslaw, K. S., Kulmala, M., Curtius, J., Baltensperger, U., Worsnop, D. R., Vehkamäki, H.
- 591 and Kirkby, J.: Molecular understanding of sulphuric acid-amine particle nucleation in the
- 592 atmosphere., Nature, 502(7471), 359–63, doi:10.1038/nature12663, 2013.
- 593 Bianchi, F., Tröstl, J., Junninen, H., Frege, C., Henne, S., Hoyle, C. R., Molteni, U.,
- Herrmann, E., Adamov, A., Bukowiecki, N., Chen, X., Duplissy, J., Gysel, M., Hutterli, M.,
- 595 Kangasluoma, J., Kontkanen, J., Kürten, A., Manninen, H. E., Münch, S., Peräkylä, O.,
- 596 Petäjä, T., Rondo, L., Williamson, C., Weingartner, E., Curtius, J., Worsnop, D. R., Kulmala,
- 597 M., Dommen, J. and Baltensperger, U.: New particle formation in the free troposphere: A
- 598 question of chemistry and timing, Science, 352 (6289), 1109-1112,
- 599 doi:10.1126/science.aad5456, 2016.
- 600 Boucher, O., Randall, D., Artaxo, P., Bretherton, C., Feingold, G., Forster, P., Kerminen, V.-
- 601 M. V.-M., Kondo, Y., Liao, H., Lohmann, U., Rasch, P., Satheesh, S. K., Sherwood, S.,
- 602 Stevens, B., Zhang, X. Y. and Zhan, X. Y.: Clouds and Aerosols. In: Climate Change 2013:
- 603 The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report
- 604 of the Intergovernmental Panel on Climate Change., 571–657,
- 605 doi:10.1017/CBO9781107415324.016, 2013.
- 606 Crounse, J. D., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G. and Wennberg, P. O.:
- 607 Autoxidation of organic compounds in the atmosphere, J. Phys. Chem. Lett., 4(20), 3513–
- 608 3520, doi:10.1021/jz4019207, 2013.
- 609 Ditchfield, R.: Self-consistent molecular-orbital methods. IX. An extended gaussian-type
- 610 basis for molecular-orbital studies of organic molecules, J. Chem. Phys., 54 (2), 724,
- 611 doi:10.1063/1.1674902, 1971.
- 612 Dunne, E. M., Gordon, H., Andreas, K., Duplissy, J., Williamson, C., Ortega, I. K., Pringle,
- 613 K. J., Adamov, A., Baltensperger, U., Barmet, P., Benduhn, F., Bianchi, F., Breitenlechner,
- 614 M., Clarke, A., Curtius, J., Dommen, J., Donahue, N. M., Ehrhart, S., Flagan, R. C., Franchin,
- 615 A., Guida, R., Hakala, J., Hansel, A., Heinritzi, M., Jokinen, T., Kangasluoma, J., Kirkby, J.,
- Kulmala, M., Kupc, A., Lawler, M. J., Lehtipalo, K., Makhmutov, V., Mathot, S., Miettinen,
- P., Nenes, A., Onnela, A., Rap, A., Reddington, C. L. S., Riccobono, F., Richards, N. A. D.,
 Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Sengupta, K., Simon, M., Sipilä,
- M., Smith, J. N., Stozkhov, Y. and Tom, A.: Global particle formation from CERN CLOUD
- 620 measurements. Science, 354 (6316) 1119-1124, doi:10.1126/science.aaf2649, 2016.

Manuscript under review for journal Atmos. Chem. Phys.

Discussion started: 29 May 2017





- 621 Duplissy, J., Merikanto, J., Franchin, A., Tsagkogeorgas, G., Kangasluoma, J., Wimmer, D.,
- 622 Vuollekoski, H., Schobesberger, S., Lehtipalo, K., Flagan, R. C., Brus, D., Donahue, N. M.,
- 623 Vehkamäki, H., Almeida, J., Amorim, A., Barmet, P., Bianchi, F., Breitenlechner, M.,
- 624 Dunne, E. M., Guida, R., Henschel, H., Junninen, H., Kirkby, J., Kürten, A., Kupc, A.,
- 625 Määttänen, A., Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Praplan, A. P.,
- 626 Riccobono, F., Rondo, L., Steiner, G., Tomé, A., Walther, H., Baltensperger, U., Carslaw, K.
- 627 S., Dommen, J., Hansel, A., Petäjä, T., Sipilä, M., Stratmann, F., Vrtala, A., Wagner, P. E.,
- 628 Worsnop, D. R., Curtius, J. and Kulmala, M.: Effect of ions on sulfuric acid-water binary
- 629 particle formation: 2. Experimental data and comparison with QC-normalized classical
- 630 nucleation theory, J. Geophys. Res. Atmos., 121 (4), 1752-1775, doi:10.1002/2015JD023538,
- 631 2016.
- 632 Ehn, M., Junninen, H., Petäjä, T., Kurtén, T., Kerminen, V.-M., Schobesberger, S.,
- 633 Manninen, H. E., Ortega, I. K., Vehkamäki, H., Kulmala, M. and Worsnop, D. R.:
- 634 Composition and temporal behavior of ambient ions in the boreal forest, Atmos. Chem.
- 635 Phys., 10(17), 8513–8530, doi:10.5194/acp-10-8513-2010, 2010.
- 636 Ehn, M., Junninen, H., Schobesberger, S., Manninen, H. E., Franchin, A., Sipilä, M., Petäjä,
- 637 T., Kerminen, V.-M., Tammet, H., Mirme, A., Mirme, S., Hõrrak, U., Kulmala, M. and
- 638 Worsnop, D. R.: An instrumental comparison of mobility and mass measurements of
- 639 atmospheric small ions, Aerosol Sci. Technol., 45(4), 522-532,
- 640 doi:10.1080/02786826.2010.547890, 2011.
- 641 Ehn, M., Kleist, E., Junninen, H., Petäjä, T., Lönn, G., Schobesberger, S., Dal Maso, M.,
- 642 Trimborn, A., Kulmala, M., Worsnop, D. R., Wahner, A., Wildt, J. and Mentel, T. F.: Gas
- 643 phase formation of extremely oxidized pinene reaction products in chamber and ambient air,
- 644 Atmos. Chem. Phys., 12(11), 5113–5127, doi:10.5194/acp-12-5113-2012, 2012.
- 645 Ehn, M., Thornton, J. a., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M.,
- 646 Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M.,
- 647 Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T.,
- 648 Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T.,
- 649 Petäjä, T., Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J. and
- 650 Mentel, T. F.: A large source of low-volatility secondary organic aerosol, Nature, 506(7489),
- 651 476–479, doi:10.1038/nature13032, 2014.
- 652 Franchin, A., Ehrhart, S., Leppä, J., Nieminen, T., Gagné, S., Schobesberger, S., Wimmer,
- 653 D., Duplissy, J., Riccobono, F., Dunne, E. M., Rondo, L., Downard, A., Bianchi, F., Kupc,
- 654 A., Tsagkogeorgas, G., Lehtipalo, K., Manninen, H. E., Almeida, J., Amorim, A., Wagner, P.
- E., Hansel, A., Kirkby, J., Kürrten, A., Donahue, N. M., Makhmutov, V., Mathot, S., 655
- 656 Metzger, A., Petäjä, T., Schnitzhofer, R., Sipilä, M., Stozhkov, Y., Tomé, A., Kerminen, V.
- 657 M., Carslaw, K., Curtius, J., Baltensperger, U. and Kulmala, M.: Experimental investigation
- 658 of ion-ion recombination under atmospheric conditions, Atmos. Chem. Phys., 15(13), 7203-
- 659 7216, doi:10.5194/acp-15-7203-2015, 2015.
- 660 Hirsikko, A., Nieminen, T., Gagné, S., Lehtipalo, K., Manninen, H. E., Ehn, M., Hõrrak, U.,
- 661 Kerminen, V.-M., Laakso, L., McMurry, P. H., Mirme, A., Mirme, S., Petäjä, T., Tammet,
- H., Vakkari, V., Vana, M. and Kulmala, M.: Atmospheric ions and nucleation: a review of 662
- 663 observations, Atmos. Chem. Phys., 11(2), 767–798, doi:10.5194/acp-11-767-2011, 2011.

Manuscript under review for journal Atmos. Chem. Phys.

Discussion started: 29 May 2017





- 664 Hoyle, C. R., Fuchs, C., Jarvinen, E., Saathoff, H., Dias, A., El Haddad, I., Gysel, M.,
- 665 Coburn, S. C., Trostl, J., Hansel, A., Bianchi, F., Breitenlechner, M., Corbin, J. C., Craven, J.,
- 666 Donahue, N. M., Duplissy, J., Ehrhart, S., Frege, C., Gordon, H., Hoppel, N., Heinritzi, M.,
- Kristensen, T. B., Molteni, U., Nichman, L., Pinterich, T., Prevôt, A. S. H., Simon, M.,
- Slowik, J. G., Steiner, G., Tome, A., Vogel, A. L., Volkamer, R., Wagner, A. C., Wagner, R.,
- 669 Wexler, A. S., Williamson, C., Winkler, P. M., Yan, C., Amorim, A., Dommen, J., Curtius,
- 670 J., Gallagher, M. W., Flagan, R. C., Hansel, A., Kirkby, J., Kulmala, M., Mohler, O.,
- 671 Stratmann, F., Worsnop, D. R. and Baltensperger, U.: Aqueous phase oxidation of sulphur
- dioxide by ozone in cloud droplets, Atmos. Chem. Phys., 16(3), 1693–1712, doi:10.5194/acp-
- 673 16-1693-2016, 2016.
- Hunter, E. P. and Lias, S. G.: Evaluated gas phase basicities and proton affinity of molecules:
- an update, J. Phys. Chem. Ref. Data, 27(3), 413–656, 1998.
- Hyttinen, N., Kupiainen-Määttä, O., Rissanen, M. P., Muuronen, M., Ehn, M. and Kurtén, T.:
- 677 Modeling the charging of highly oxidized cyclohexene ozonolysis products using nitrate-
- 678 based chemical ionization, J. Phys. Chem. A, 119(24), 6339–6345,
- 679 doi:10.1021/acs.jpca.5b01818, 2015.
- 680 Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné,
- 681 S., Ickes, L., Kürten, A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S.,
- Tsagkogeorgas, G., Wimmer, D., Amorim, A., Bianchi, F., Breitenlechner, M., David, A.,
- 683 Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud,
- W., Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy,
- E. R., Makhmutov, V., Mathot, S., Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T.,
- Onnela, A., Pereira, P., Petäjä, T., Schnitzhofer, R., Seinfeld, J. H., Sipilä, M., Stozhkov, Y.,
- 687 Stratmann, F., Tomé, A., Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H.,
- Weingartner, E., Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U.
- 689 and Kulmala, M.: Role of sulphuric acid, ammonia and galactic cosmic rays in atmospheric
- 690 aerosol nucleation., Nature, 476(7361), 429–33, doi:10.1038/nature10343, 2011.
- 691 Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M.,
- 692 Simon, M., Yan, C., Almeida, J., Tröstl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov,
- 693 A., Amorim, A., Bernhammer, A.-K., Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X.,
- 694 Craven, J., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Hakala, J.,
- 695 Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim, J., Krapf, M., Kürten, A.,
- Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Molteni, U., Onnela, A., Peräkylä,
- 697 O., Piel, F., Petäjä, T., Praplan, A. P., Pringle, K., Rap, A., Richards, N. A. D., Riipinen, I.,
- 698 Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger, S., Scott, C. E., Seinfeld, J. H.,
- 699 Sipilä, M., Steiner, G., Stozhkov, Y., Stratmann, F., Tomé, A., Virtanen, A., Vogel, A. L.,
- 700 Wagner, A. C., Wagner, P. E., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Zhang,
- 700 Wagner, A. C., Wagner, T. E., Weingardier, E., Winnier, D., Whiker, T. M., Te, T., Zhang, 701 X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R., Baltensperger, U., Kulmala,
- 702 M., Carslaw, K. S. and Curtius, J.: Ion-induced nucleation of pure biogenic particles, Nature,
- 703 533(7604), 521–526, doi:10.1038/nature17953, 2016.
- Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili,
- 705 W. and McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: A
- 706 review of observations, J. Aerosol Sci., 35(2), 143–176, doi:10.1016/j.jaerosci.2003.10.003,
- 707 2004.

Manuscript under review for journal Atmos. Chem. Phys.

Discussion started: 29 May 2017





- 708 Kürten, A., Bianchi, F., Almeida, J., Kupiainen-Määttä, O., Dunne, E. M., Duplissy, J.,
- Williamson, C., Barmet, P., Breitenlechner, M., Dommen, J., Donahue, N. M., Flagan, R. C.,
- 710 Franchin, A., Gordon, H., Hakala, J., Hansel, A., Heinritzi, M., Ickes, L., Jokinen, T.,
- Kangasluoma, J., Kim, J., Kirkby, J., Kupc, A., Lehtipalo, K., Leiminger, M., Makhmutov,
- V., Onnela, A., Ortega, I. K., Petäjä, T., Praplan, A. P., Riccobono, F., Rissanen, M. P.,
- Rondo, L., Schnitzhofer, R., Schobesberger, S., Smith, J. N., Steiner, G., Stozkhov, Y.,
- 714 Tomé, A., Tröstl, J., Tsagkogeorgas, G., Wagner, P. E., Wimmer, D., Ye, P., Baltensperger,
- 715 U., Carslaw, K. S., Kulmala, M. and Curtius, J.: Experimental particle formation rates
- spanning tropospheric sulfuric acid and ammonia abundances, ion production rates, and
- 717 temperatures, J. Geophys. Res. Atmos., 121, doi:10.1002/2015JD023908, 2016.
- 718 Lehtipalo, K., Rondo, L., Kontkanen, J., Schobesberger, S., Jokinen, T., Sarnela, N., Kürten,
- A., Ehrhart, S., Franchin, A., Nieminen, T., Riccobono, F., Sipilä, M., Yli-Juuti, T., Duplissy,
- 720 J., Adamov, A., Ahlm, L., Almeida, J., Amorim, A., Bianchi, F., Breitenlechner, M.,
- 721 Dommen, J., Downard, A. J., Dunne, E. M., Flagan, R. C., Guida, R., Hakala, J., Hansel, A.,
- Jud, W., Kangasluoma, J., Kerminen, V.-M., Keskinen, H., Kim, J., Kirkby, J., Kupc, A.,
- 723 Kupiainen-Määttä, O., Laaksonen, A., Lawler, M. J., Leiminger, M., Mathot, S., Olenius, T.,
- 724 Ortega, I. K., Onnela, A., Petäjä, T., Praplan, A. P., Rissanen, M. P., Ruuskanen, T. M.,
- 725 Santos, F. D., Schallhart, S., Schnitzhofer, R., Simon, M., Smith, J. N., Tröstl, J.,
- 726 Tsagkogeorgas, G., Tomé, A., Vaattovaara, P., Vehkämäki, H., Vrtala, A. E., Wagner, P. E.,
- Williamson, C., Wimmer, D., Winkler, P. M., Virtanen, A., Donahue, N. M., Carslaw, K. S.,
- 728 Baltensperger, U., Riipinen, I., Curtius, J., Worsnop, D. R. and Kulmala, M.: The effect of
- 729 acid-base clustering and ions on the growth of atmospheric nano-particles, Nat. Commun.,7,
- 730 doi:10.1038/ncomms11594, 2016.
- 731 Merikanto, J., Spracklen, D. V, Mann, G. W., Pickering, S. J. and Carslaw, K. S.: Impact of
- nucleation on global CCN, Atmos. Chem. Phys., 8601–8616, 2009.
- 733 Merikanto, J., Duplissy, J., Määttänen, A., Henschel, H., Donahue, N. M., Brus, D.,
- 734 Schobesberger, S., Kulmala, M. and Vehkamäki, H.: Effect of ions on sulfuric acid-water
- 735 binary particle formation I: Theory for kinetic and nucleation-type particle formation and
- 736 atmospheric implications, J. Geophys. Res. Atmos., 121, 1736–1751,
- 737 doi:10.1002/2015JD023538, 2016.
- 738 Metzger, A., Verheggen, B., Dommen, J., Duplissy, J., Prevot, A. S. H., Weingartner, E.,
- 739 Riipinen, I., Kulmala, M., Spracklen, D. V, Carslaw, K. S. and Baltensperger, U.: Evidence
- 740 for the role of organics in aerosol particle formation under atmospheric conditions., Proc.
- 741 Natl. Acad. Sci. U. S. A., 107(15), 6646–51, doi:10.1073/pnas.0911330107, 2010.
- 742 Nadykto, A. B.: Uptake of neutral polar vapor molecules by charged clusters/particles:
- Enhancement due to dipole-charge interaction, J. Geophys. Res. Atmos., 108(D23), 4717,
- 744 doi:10.1029/2003JD003664, 2003.
- 745 Ortega, I. K., Kupiainen, O., Kurtén, T., Olenius, T., Wilkman, O., McGrath, M. J.,
- 746 Loukonen, V. and Vehkamäki, H.: From quantum chemical formation free energies to
- 747 evaporation rates, Atmos. Chem. Phys., 12(1), 225–235, doi:10.5194/acp-12-225-2012, 2012.
- 748 Praplan, A. P., Schobesberger, S., Bianchi, F., Rissanen, M. P., Ehn, M., Jokinen, T.,
- Junninen, H., Adamov, A., Amorim, A., Dommen, J., Duplissy, J., Hakala, J., Hansel, A.,
- 750 Heinritzi, M., Kangasluoma, J., Kirkby, J., Krapf, M., Kürten, A., Lehtipalo, K., Riccobono,

Manuscript under review for journal Atmos. Chem. Phys.

Discussion started: 29 May 2017





- 751 F., Rondo, L., Sarnela, N., Simon, M., Tomé, A., Tröstl, J., Winkler, P. M., Williamson, C.,
- 752 Ye, P., Curtius, J., Baltensperger, U., Donahue, N. M., Kulmala, M. and Worsnop, D. R.:
- 753 Elemental composition and clustering behaviour of α-pinene oxidation products for different
- 754 oxidation conditions, Atmos. Chem. Phys., 15(8), 4145–4159, doi:10.5194/acp-15-4145-
- 755 2015, 2015.
- 756 Riccobono, F., Schobesberger, S., Scott, C. E., Dommen, J., Ortega, I. K., Rondo, L.,
- 757 Almeida, J., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Downard, A., Dunne, E.
- 758 M., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Hansel, A., Junninen, H., Kajos, M.,
- 759 Keskinen, H., Kupc, A., Kürten, A., Kvashin, A. N., Laaksonen, A., Lehtipalo, K.,
- Makhmutov, V., Mathot, S., Nieminen, T., Onnela, A., Petäjä, T., Praplan, A. P., Santos, F.
- 761 D., Schallhart, S., Seinfeld, J. H., Sipilä, M., Spracklen, D. V, Stozhkov, Y., Stratmann, F.,
- 762 Tomé, A., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y., Vrtala, A., Wagner, P. E.,
- Donahue, N. M., Kirkby, J., Kulmala, M., Worsnop, D. R. and Baltensperger, U.: Oxidation
- products of biogenic emissions contribute to nucleation of atmospheric particles, Science,
- 765 344, 717–721, doi:10.1126/science.1243527, 2014.
- 766 Rissanen, M. P., Kurtén, T., Sipilä, M., Thornton, J. A., Kausiala, O., Garmash, O.,
- 767 Kjaergaard, H. G., Petäjä, T., Worsnop, D. R., Ehn, M. and Kulmala, M.: Effects of chemical
- 768 complexity on the autoxidation mechanisms of endocyclic alkene ozonolysis products: From
- methylcyclohexenes toward understanding α-pinene, J. Phys. Chem. A, 119(19), 4633–4650,
- 770 doi:10.1021/jp510966g, 2015.
- 771 Schobesberger, S., Junninen, H., Bianchi, F., Lönn, G., Ehn, M., Lehtipalo, K., Dommen, J.,
- 772 Ehrhart, S., Ortega, I. K., Franchin, A., Nieminen, T., Riccobono, F., Hutterli, M., Duplissy,
- 773 J., Almeida, J., Amorim, A., Breitenlechner, M., Downard, A. J., Dunne, E. M., Flagan, R.
- 774 C., Kajos, M., Keskinen, H., Kirkby, J., Kupc, A., Kürten, A., Kurtén, T., Laaksonen, A.,
- 775 Mathot, S., Onnela, A., Praplan, A. P., Rondo, L., Santos, F. D., Schallhart, S., Schnitzhofer,
- R., Sipilä, M., Tomé, A., Tsagkogeorgas, G., Vehkamäki, H., Wimmer, D., Baltensperger, U.,
- 777 Carslaw, K. S., Curtius, J., Hansel, A., Petäjä, T., Kulmala, M., Donahue, N. M. and
- 778 Worsnop, D. R.: Molecular understanding of atmospheric particle formation from sulfuric
- 779 acid and large oxidized organic molecules., Proc. Natl. Acad. Sci. U. S. A., 110(43), 17223-
- 780 8, doi:10.1073/pnas.1306973110, 2013.
- 781 Shuman, N. S., Hunton, D. E. and Viggiano, A. A.: Ambient and modified atmospheric ion
- 782 chemistry: from top to bottom, Chem. Rev., 115(10), 4542–4570, doi:10.1021/cr5003479,
- 783 2015.
- 784 Steiner, G., Jokinen, T., Junninen, H., Sipilä, M., Petäjä, T., Worsnop, D., Reischl, G. P. and
- 785 Kulmala, M.: High-resolution mobility and mass spectrometry of negative ions produced in a
- 786 241 Am aerosol charger, Aerosol Sci. Technol., 48 (3), 261–270,
- 787 doi:10.1080/02786826.2013.870327, 2014.
- 788 Tröstl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege,
- 789 C., Bianchi, F., Wagner, R., Simon, M., Lehtipalo, K., Williamson, C., Craven, J. S.,
- 790 Duplissy, J., Adamov, A., Almeida, J., Bernhammer, A.-K., Breitenlechner, M., Brilke, S.,
- Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs, C., Guida, R., Gysel, M., Hansel, A.,
- Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Keskinen, H., Kim, J., Krapf, M., Kürten, A., Laaksonen, A., Lawler, M. J., Leiminger, M., Mathot, S., Möhler, O., Nieminen,
- 794 T., Onnela, A., Petäjä, T., Piel, F., Miettinen, P., Rissanen, M. P., Rondo, L., Sarnela, N.,

Discussion started: 29 May 2017

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- 795 Schobesberger, S., Sengupta, K., Sipilä, M., Smith, J. N., Steiner, G., Tomé, A., Virtanen, A.,
- Wagner, A. C., Weingartner, E., Wimmer, D., Winkler, P. M., Ye, P., Carslaw, K. S.,
- 797 Curtius, J., Dommen, J., Kirkby, J., Kulmala, M., Riipinen, I., Worsnop, D. R., Donahue, N.
- 798 M. and Baltensperger, U.: The role of low-volatility organic compounds for initial particle
- 799 growth in the atmosphere, Nature, 533, 527–531, doi:doi:10.1038/nature18271, 2016.
- 800 Zhao, Y. and Truhlar, D. G.: The M06 suite of density functionals for main group
- 801 thermochemistry, thermochemical kinetics, noncovalent interactions, excited states, and
- 802 transition elements: Two new functionals and systematic testing of four M06-class
- 803 functionals and 12 other functionals, Theor. Chem. Acc., 120(1-3), 215-241,
- 804 doi:10.1007/s00214-007-0310-x, 2008.