

1 **Liquid-liquid phase separation in particles containing secondary organic**
2 **material free of inorganic salts**

3

4 Mijung Song^{1,2}, Pengfei Liu³, Scot T. Martin^{3,4}, Allan K. Bertram^{2*}

5 [1] {Department of Earth and Environmental Sciences, Chonbuk National University, Jeollabuk-
6 do, Republic of Korea}

7 [2] {Department of Chemistry, University of British Columbia, Vancouver, BC, V6T 1Z1, Canada}

8 [3] {John A. Paulson School of Engineering and Applied Sciences, Harvard University,
9 Cambridge, Massachusetts 02138, USA}

10 [4] {Department of Earth and Planetary Sciences, Harvard University, Cambridge, Massachusetts
11 02138, USA}

12 Correspondence to: A. K. Bertram (bertram@chem.ubc.ca)

13

14 **Abstract**

15 Particles containing secondary organic material (SOM) are ubiquitous in the atmosphere and play
16 a role in climate and air quality. Recently, research has shown that liquid-liquid phase separation
17 (LLPS) occurs at high relative humidities (RH) (greater than ~ 95 %) in α -pinene-derived SOM
18 particles free of inorganic salts while LLPS does not occur in isoprene-derived SOM particles free
19 of inorganic salts. We expand on these findings by investigating LLPS at 290 ± 1 K in SOM
20 particles free of inorganic salts produced from ozonolysis of β -caryophyllene, ozonolysis of
21 limonene, and photo-oxidation of toluene. LLPS was observed at greater than ~95 % RH in the
22 biogenic SOM particles derived from β -caryophyllene and limonene while LLPS was not observed
23 in the anthropogenic SOM particles derived from toluene. This work combined with the earlier
24 work on LLPS in SOM particles free of inorganic salts suggests that the occurrence of LLPS in
25 SOM particles free of inorganic salts is related to the oxygen-to-carbon elemental ratio (O:C) of
26 the organic material. These results help explain the difference between the hygroscopic parameter

1 κ of SOM particles measured above and below water saturation in the laboratory and field, and
2 have implications for predicting the cloud condensation nucleation properties of SOM particles.

3

4 **1 Introduction**

5 Secondary organic material (SOM) is produced in the atmosphere by the oxidation of volatile
6 organic compounds (VOCs) such as α -pinene and isoprene from trees and toluene from
7 anthropogenic sources. Once formed, the low volatility and semivolatile oxidation products can
8 partition to the particle phase to form SOM containing particles (Hallquist et al., 2009; Ervens et
9 al., 2011). SOM accounts for approximately 20 – 80 % of the submicrometer particle mass in the
10 atmosphere (Zhang et al., 2007; Jimenez et al., 2009). Although the exact chemical composition
11 of SOM in atmospheric particles remains an active area of research, laboratory and field studies
12 have shown that the average oxygen-to-carbon elemental ratio (O:C) of these particles ranges from
13 0.2 to 1.0 (Chen et al., 2009; Jimenez et al., 2009; Heald et al., 2010; Takahama et al., 2011).

14 As the relative humidity (RH) varies in the atmosphere, SOM containing particles can undergo
15 several different phase transitions with implications for the cloud condensation nuclei (CCN)
16 properties, optical properties, reactivity, and growth of these particles (Martin et al., 2000;
17 Raymond and Pandis, 2002; Bilde and Svenningsson, 2004; Zuend et al., 2010; Kuwata and Martin,
18 2012; Brunamonti et al., 2015). One possible phase transition that SOM particles may undergo as
19 RH varies in the atmosphere is liquid-liquid phase separation (LLPS) (Pankow, 2003; Marcolli et
20 al., 2006; Ciobanu et al., 2009; Zuend and Seinfeld, 2012; Veghte et al., 2013; O'Brien et al.,
21 2015). LLPS in particles containing both SOM and inorganic salts has been the focus of many
22 recent studies. These studies have shown that SOM particles mixed with inorganic salts can
23 undergo LLPS in the atmosphere when the O:C of the organic material is less than 0.56, but LLPS
24 may not occur when the O:C of the organic material is greater than 0.80 (Bertram et al., 2011;
25 Krieger et al., 2012; Smith et al., 2012; Song et al., 2012a; Schill and Tolbert, 2013; Song et al.,
26 2013; You et al., 2013; You et al., 2014).

27 Recently, researchers have also focused on LLPS in SOM particles free of inorganic salts. Petters
28 et al. (2006) suggested that a miscibility gap in particles containing organic polymers at high RH
29 may lead to a non-classical pathway for CCN activation. Renbaum-Wolff et al. (2016) showed that
30 α -pinene-derived SOM free of inorganic salts can undergo LLPS at high RH values (~95 to 100 %),

1 **which could result in** altered CCN properties. In addition, they showed that LLPS in SOM particles
2 will lead to a different hygroscopic parameter, κ , at subsaturated conditions compared to
3 supersaturated conditions. The implication is that the CCN activity of SOM particles, if they
4 undergo LLPS, is higher than predicted from subsaturated hygroscopicity measurements. **Related,**
5 **Hodas et al. (2016), using a combination of measurements and modelling of surrogates of**
6 **oligomers, showed that the prevalence of LLPS at high RH can contribute to differences in**
7 **hygroscopicity above and below water saturation.** Most recently, Rastak et al. (2017) observed that
8 isoprene-derived SOM particles do not undergo LLPS even at high RH. Rastak et al. (2017) used
9 these results together with thermodynamic calculations to explain the hygroscopic properties of
10 biogenic organic aerosol particles in the laboratory and the field.

11 Here we expand on the studies by Renbaum-Wolff et al. (2016) and Rastak et al. (2017) by
12 investigated LLPS in SOM particles generated by the ozonolysis of limonene, ozonolysis of β -
13 caryophyllene, and photo-oxidation of toluene. Limonene and β -caryophyllene are both biogenic
14 VOCs, while toluene is an anthropogenic VOC (Kanakidou et al., 2005). Both limonene-derived
15 and β -caryophyllene-derived SOM particles have been used as proxies of biogenic SOM particles
16 in the atmosphere (Bateman et al., 2009; Alfarra et al., 2012; Kundu et al., 2012; Frosch et al.,
17 2013; Liu et al., 2013), while toluene-derived SOM has been used as a proxy for anthropogenic
18 SOM particles (Pandis et al., 1992; Robinson et al., 2013; Liu et al., 2016; Song et al., 2016; Ye
19 et al., 2016).

20

21 **2 Methods**

22 **2.1 Production of secondary organic materials**

23 SOM particles were generated via β -caryophyllene ozonolysis and limonene ozonolysis in a flow
24 tube reactor (Table 1) and via toluene photo-oxidation in an oxidation flow reactor (OFR) (Table
25 2). The method of particle generation in the flow tube reactor was described in Shrestha et al.
26 (2013) and the methods of particle generation in the OFR (Kang et al., 2007) were given in Liu et
27 al. (2015). The flow tube reactor was operated at a flow rate of 3.5 L min^{-1} (with a residence time
28 of 38 s) and $< 5\%$ RH. The OFR was operated at flow rates of 7.0 to 9.5 L min^{-1} (with a residence
29 time of 80 to 110 s) and $13 \pm 3\%$ RH. Both reactors were operated at a temperature of $293 \pm 2 \text{ K}$.

1 Table 1 lists the experimental conditions for the production of SOM via ozonolysis. For the particle
2 generation via ozonolysis, ozone was produced by irradiating pure air (Aadco 737 Pure Air
3 Generator) with ultraviolet emission from a mercury lamp ($\lambda = 185$ nm). Ozone concentrations
4 used for ozonolysis ranged from 12 - 30 ppm for β -caryophyllene and 13 - 30 ppm for limonene
5 (Table 1). β -caryophyllene and limonene (Sigma Aldrich, ≥ 99 %) were dissolved in 2-butanol
6 (Sigma-Aldrich, ≥ 99.5 %). These organic solutions were injected into a glass round-bottom flask
7 held at 310 K, where the organic liquids vaporized at the tip of a syringe. The organic vapor was
8 then swept into the reactor where ozonolysis took place to form SOM and particles. The injected
9 precursor concentrations were 0.03 - 0.7 ppm for β -caryophyllene and 0.07 - 2.0 ppm for limonene
10 in the main flow of the reactor. In the ozonolysis experiments, butanol served as an OH radical
11 scavenger.

12 Table 2 presents the experimental conditions for the production of SOM via photo-oxidation. For
13 the particle generation via photo-oxidation, hydroxyl radicals were produced in the OFR by the
14 photochemical reactions:



17

18 Ozone was again produced by irradiating pure air (Aadco 737 Pure Air Generator) with ultraviolet
19 emission from a mercury lamp ($\lambda = 185$ nm). Ozone concentrations used in the photo-oxidation
20 studies were 30 ppm (Table 2). Toluene (Sigma-Aldrich, 99 %) was injected and vaporized in a
21 flask, and the vapors were swept into the OFR by purified air. The injected toluene concentrations
22 were 0.2 - 1.0 ppm.

23 The mass concentration of SOM particles during the generation process was determined from
24 measurements of the number-diameter distribution of SOM particles in the flow tube reactor or
25 OFR and assuming a material density of 1200 kg m^{-3} (Liu et al., 2013). The number-diameter
26 distributions were measured with a scanning mobility particle sizer (SMPS; TSI Inc.). The O:C
27 ratio of the toluene-derived SOM studied here was determined using a high resolution aerosol mass

1 spectrometer (HR-ToF-AMS; Aerodyne Research Inc.). Data analysis was based on the approach
2 described by Chen et al. (2011).

3

4 **2.2 Production of supermicron SOM particles on hydrophobic substrates**

5 At the outlet of the flow tube reactor and OFR, the sub-micrometer SOM particles were collected
6 on hydrophobic surfaces. The limonene-derived SOM and toluene-derived SOM particles were
7 collected onto glass slides coated with trichloro(1*H*,1*H*,2*H*,2*H*-perfluorooctyl)silane (Sigma-
8 Aldrich, 97%). The coating procedure is described in Knopf (2003). The β -caryophyllene-derived
9 SOM particles were collected onto Teflon substrates.

10 Two different methods were used to collect submicron particles on hydrophobic substrates (see
11 Tables 1 and 2). The first method used was an electrostatic precipitator (TSI 3089, USA). In this
12 case, the resulting SOM particles on the hydrophobic substrates were smaller than $\sim 10 \mu\text{m}$. **From**
13 **experience in our laboratory, detection of LLPS with our microscope setup is the clearest when**
14 **the size of the particles are roughly 20 - 80 μm . As a result,** the following method was used to
15 coagulate the sub-10 μm particles into 20 - 80 μm particles: first the substrate containing the SOM
16 particles was placed in a RH-controlled flow-cell (Parsons et al., 2004; Pant et al., 2006; Song et
17 al., 2012b). The RH in the flow-cell was then set to over 100 % for 30 - 60 min to grow and
18 coagulate the SOM particles. The RH in the flow-cell was then decreased to $\sim 80 - 90$ % RH to
19 evaporate the water. During the experiments, the particles were observed using a reflectance
20 microscope (Zeiss Axiotech, 50 \times). These growth and coagulation processes resulted in SOM
21 particles consisting of 20 - 80 μm in diameter (Song et al., 2015; Renbaum-Wolff et al., 2016).

22 In the second method used to collect SOM particles collected on a hydrophobic substrate, a single
23 stage impactor was used (Prenni et al., 2009; Pöschl et al., 2010; Hosny et al., 2016). In this case,
24 the SOM particles after collection were as big as 100 μm due to coagulation during the collection
25 process. Since the particles were already large enough for the LLPS experiments, they were used
26 directly without the need for the growth and coagulation experiments described above. Both
27 methods used to collect SOM particles collected both the water-soluble and water-insoluble
28 components of the SOM particles.

29

1 **2.3 Optical microscopy of supermicron SOM particles**

2 For the LLPS experiments, the hydrophobic substrate containing SOM particles with sizes in the
3 range of 20 to 80 μm in diameters was mounted in a temperature and RH controlled flow-cell
4 coupled to an optical reflectance microscope (Zess Axiotech, 50 \times objective) (Parsons et al., 2004;
5 Pant et al., 2006; Song et al., 2012b). The temperature of the cell was 290 ± 1 K in all experiments.
6 RH in the cell was regulated by varying the ratio of a dry and humidified N_2 flow. The total flow
7 rate was ~ 1200 sccm. **The RH was determined from measurements of the temperature with a**
8 **thermocouple and measurements of the dew point/frost point with a chilled mirror sensor (General**
9 **Eastern, Canada). The RH was calibrated using the deliquescence RH for pure ammonium sulfate**
10 **particles (80% RH at 293 K, Martin, 2000). After calibration, the uncertainty of the RH was ± 2.0 %**
11 **based on the reproducibility of multiple deliquesce measurements.** At the beginning of LLPS
12 experiments the SOM particles were equilibrated at ~ 100 % RH for 15 minutes. Then the RH was
13 reduced from ~ 100 to ~ 0 % RH at a rate of 0.1 to 0.5 % RH min^{-1} , and subsequently increased to
14 ~ 100 % RH at a rate of 0.1 to 0.5 % RH min^{-1} . **We did not observe a dependence of LLPS on the**
15 **RH ramp rate, although only a narrow range of rates were used.** During the humidity cycle, optical
16 images of the SOM particles were recorded every 5 - 10 seconds using a CCD camera.

17

18 **3 Results and Discussion**

19 **3.1 β -caryophyllene-derived and limonene-derived SOM particles**

20 Humidity cycles at 290 ± 1 K were performed for β -caryophyllene-derived SOM particles
21 generated with mass concentrations of 15 - 4000 $\mu\text{g m}^{-3}$ and limonene-derived SOM generated
22 with mass concentrations of 80 - 7000 $\mu\text{g m}^{-3}$ (Table 1). In all cases, LLPS was observed at high
23 RH. Table 1 summarizes the results during humidity cycles. Shown in Fig. 1a and Movie S1
24 (Supplementary Material) are examples of optical images of a β -caryophyllene-derived SOM
25 particle as a function of increasing RH for the particle mass concentrations of 2000 - 4000 $\mu\text{g m}^{-3}$.
26 Shown in Fig. 1b and Movie S2 (Supplementary Material) are examples of optical images of a
27 limonene-derived SOM particle as a function of increasing RH for the particle mass concentrations
28 of 7000 $\mu\text{g m}^{-3}$. For both types of SOM particles, only one phase was observed for RH values from
29 0 to ~ 90 %. Note, the light-colored circle in the center of the particles at 90.5 % RH for β -

1 caryophyllene-derived SOM and at 95.0 % RH for limonene-derived SOM is an optical effect due
2 to the light scattering from a hemispherical particle (Bertram et al., 2011). In Fig. 1, LLPS is
3 observed at 91.5 % RH for the β -caryophyllene-derived SOM particle and at 95.3 % RH for the
4 limonene-derived SOM particle. LLPS began with the formation of many small inclusions of a
5 second phase, and in both cases the phase transition occurred over a narrow range of RH. The
6 small inclusions coagulated to larger droplets in the β -caryophyllene-derived SOM at 92.5 % RH
7 and in the limonene-derived SOM at 96.1 % RH (Figs. 1a and 1b). **At the highest RH investigated,**
8 **a core-shell morphology is observed. Such a core-shell morphology on a hydrophobic**
9 **substrate has been observed previously in organic/ammonium sulfate/H₂O particles by Song et al.**
10 **(2012b), although a different morphology can result in the absence of the hydrophobic substrate**
11 **(Reid et al., 2011).** After formation of the core-shell morphology consisting of an inner and outer
12 phase, the two liquid phases co-existed as high as ~100 % RH. **We assume that the inner phase is**
13 **a water-rich phase while the other phase is an organic-rich phase, since the size of the inner phase**
14 **decreases as the RH decreases (Renbaum-Wolff et al., 2016).** The surface tension of water and the
15 **surface tension of more-oxidized and less oxidized organics is consistent with this assumption**
16 **(Jasper, 1972).** Upon drying, the two liquid phases merge into one liquid phase. This **mixing**
17 process occurred at 90.9 % RH for β -caryophyllene-derived SOM and 95.6 % RH for limonene-
18 derived SOM. Movies of the **mixing** process are shown in the Supplementary Material (Movies
19 S3 and S4).

20 **Shilling et al. (2009) showed that the O:C of SOM can depend on the particle mass concentration**
21 **used to generate the SOM.** To determine whether the occurrence of LLPS depends on the SOM
22 particle mass concentrations used when generating the SOM, particle mass concentrations ranging
23 from 15 – 7000 $\mu\text{g m}^{-3}$ were investigated (Table 1). Illustrated in Fig. 2a and 2b is the RH at which
24 two phases were observed during humidity cycles as a function of the mass concentrations of the
25 β -caryophyllene-derived and limonene-derived SOM samples. Triangles represent **mixing** relative
26 humidities (MRH) of two liquid phases upon drying and circles represents separation relative
27 humidity (SRH) upon moistening.

28 LLPS was observed at 93.6 ± 1.5 % RH in the β -caryophyllene-derived SOM particles for the
29 particle mass concentrations of 15 – 4000 $\mu\text{g m}^{-3}$ (Fig. 2a). In the limonene-derived SOM particles,
30 LLPS occurred at 96.1 ± 2.1 % RH for the particle mass concentrations of 80 – 7000 $\mu\text{g m}^{-3}$ (Fig.

1 2b). LLPS occurred at 96.0 ± 0.7 % RH in α -pinene-derived SOM particles for the mass
2 concentrations of $75 - 11000 \mu\text{g m}^{-3}$ (Renbaum-Wolff et al., 2016) (Fig. 2c). As shown in Fig. 2,
3 the SRH and MRH of the β -caryophyllene-derived SOM, limonene-derived SOM, and α -pinene-
4 derived SOM particles do not depend strongly on the SOM particle mass concentrations used to
5 generate the SOM.

6

7 **3.2 Toluene-derived SOM**

8 Humidity cycles were also performed for SOM particles generated from photo-oxidation of
9 toluene using particle mass concentrations of $80 - 1000 \mu\text{g m}^{-3}$ in the reactor (Table 2). None of
10 the toluene-derived SOM particles underwent LLPS during RH cycling even at high RH (Table 2).

11 Shown in Fig. 3 and Movie S5 (Supplementary Material) are optical images of a toluene-derived
12 SOM particle for mass concentrations of $80 - 100 \mu\text{g m}^{-3}$. Images in Fig. 3 and Movie S5 were
13 recorded as the RH was increased. No LLPS was observed in the SOM particle during RH cycling
14 between 0 and 100 %. Rastak et al. (2017) did not observe LLPS in isoprene-derived SOM particles
15 for the mass concentrations of $60 - 1000 \mu\text{g m}^{-3}$

16

17 **3.3 Relation between LLPS and O:C.**

18 Summarized in Table 3 are the average SRH and MRH values determined in our work and by
19 Renbaum-Wolff et al. (2016) and Rastak et al. (2017). **The average SRH and MRH values are
20 based on several cycles of RH for different SOM mass concentrations. Also included in Table 3
21 are the average O:C values for the studied SOM particles. The average O:C values for SOM from
22 toluene photo-oxidation are based on the current study (Section 2.1). The average O:C values for
23 the other SOM are taken from the literature. Since the average O:C can depend on oxidant time
24 and oxidation conditions, we chose literature data that were closest to the experimental conditions
25 used when studying LLPS (Supplementary Material, Section S1 and Table S1).**

26 Based on the data shown in Table 3, there appears to be a relationship between the occurrence of
27 LLPS and the average O:C of the organic material: **when the average O:C was between 0.34 and
28 0.44, LLPS was observed, but when the average O:C was between 0.52 and 1.30, LLPS was not**

1 observed. This trend is also apparent in Fig. 4a, where SRH data in Table 3 is plotted versus the
2 O:C data in Table 3.

3 SOM have an average O:C and a spread (or distribution) of O:C values. Similar to SOM, systems
4 containing two organics and water also have a spread in O:C and an average O:C. Hence, as a
5 starting point to understanding LLPS in SOM, we considered previous studies that explored the
6 miscibility gap in bulk solutions containing two organics and water (see Table 1 in Ganbavale et
7 al., 2015). When the average O:C of the organic material was low in a system containing two
8 organic components with water, LLPS was observed. For example, LLPS was observed in a
9 mixture of 1-butanol (O:C = 0.25), 1-propanol (O:C = 0.20), and water (Gomis-Yagües et al., 1998)
10 and in a mixture of 1-pentanol (O:C = 0.20), acetone (O:C=0.33), and water (Tiryaki et al., 1994).
11 On the other hand, when the average O:C of the organic material was high in a system containing
12 two organics and water, LLPS was not observed. For example, LLPS was not observed in a mixture
13 of acetic acid (O:C=1.00), ethanol (O:C=0.50), and water (Pickering, 1893). We conclude that the
14 relationship between average O:C and LLPS in SOM observed here is not inconsistent with
15 previous LLPS studies in systems containing two organics and water. In addition to the average
16 O:C, the spread in O:C in organic mixtures will also be important for LLPS (Renbaum-Wolff et
17 al., 2016).

18

19 **4. Implications**

20 As mentioned in the introduction, Petters et al. (2006), Hodas et al. (2016), Renbaum-Wolff et al.
21 (2016), and Rastak et al. (2017) showed using thermodynamic calculations that SOM particles that
22 undergo LLPS at high RH values have modified CCN properties. Hence, LLPS should be
23 considered when predicting the CCN properties of SOM particles derived from α -pinene
24 ozonolysis, β -caryophyllene ozonolysis, and limonene ozonolysis. A caveat is that the mass
25 concentrations used when generating the SOM particles in our experiments was larger than
26 normally found in the atmosphere (Zhang et al., 2007; Jimenez et al., 2009; Spracklen et al., 2011;
27 Li et al., 2015). Additional studies are needed to confirm LLPS in SOM particles generated using
28 more atmospherically relevant SOM mass concentrations.

1 Discrepancy between the hygroscopic parameter, κ , (Petters and Kreidenweis, 2007) measured
2 below water saturation (κ_{HGF}) and above water saturation (κ_{CCN}) in SOM particles have been
3 reported in several studies (Petters et al., 2006; Prenni et al., 2007; Juranyi et al., 2009; Petters et
4 al., 2009; Good et al., 2010; Irwin et al., 2010; Massoli et al., 2010; Dusek et al., 2011; Irwin et
5 al., 2011; Hersey et al., 2013; Pajunoja et al., 2015; Zhao et al., 2016). Petters et al. (2006), Hodas
6 et al. (2016), Renbaum-Wolff et al. (2016) and Rastak et al. (2017) suggested that such
7 discrepancies are expected in systems that undergo LLPS at high RH. Summarized in Table 4 and
8 Fig. 4b is literature data on the difference between κ_{HGF} and κ_{CCN} (denoted $\Delta\kappa$) as a function of
9 average O:C of the organic material studied (Prenni et al., 2007; Massoli et al., 2010; Pajunoja et
10 al., 2015). The experimental conditions (oxidation time and oxidation level) for the studies
11 reported in Table 4 were not necessarily similar to the experimental conditions for the studies
12 reported in Table 3, even if the same precursor volatile organic compound was used.

13 Figure 4b suggests that $\Delta\kappa$ is related to the average O:C of the organic material. Figure 4a and 4b
14 combined suggests that when the average O:C is small, LLPS occurs and the difference between
15 κ_{HGF} and κ_{CCN} is large. On the other hand, when the average O:C is large, LLPS does not occur
16 and the difference between κ_{HGF} and κ_{CCN} is small. Figure 4 provides additional support for the
17 suggestion that the LLPS is related to the discrepancies between κ_{HGF} and κ_{CCN} .

18

19 **Acknowledgments**

20 This work was supported by the Natural Sciences and Engineering Research Council of Canada.
21 Support from the US National Science Foundation (AGS-1640378) and the US Department of
22 Energy (DE-SC0012792) is also acknowledged. Mijung Song acknowledges support from a
23 National Research Foundation of Korea (NRF) grant funded by the Korea Government (MSIP)
24 (2016R1C1B1009243).

25

26 **References**

27 Aiken, A. C., Decarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S., Ulbrich,
28 I. M., Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q., Trimborn, A., Northway, M.,

1 Ziemann, P. J., Canagaratna, M. R., Onasch, T. B., Alfarra, M. R., Prevot, A. S. H., Dommen,
2 J., Duplissy, J., Metzger, A., Baltensperger, U., and Jimenez, J. L.: O/C and OM/OC ratios of
3 primary, secondary, and ambient organic aerosols with high-resolution time-of-flight aerosol
4 mass spectrometry, *Environ. Sci. Technol.*, 42, 4478-4485, Doi 10.1021/Es703009q, 2008.

5 Alfarra, M. R., Hamilton, J. F., Wyche, K. P., Good, N., Ward, M. W., Carr, T., Barley, M. H.,
6 Monks, P. S., Jenkin, M. E., Lewis, A. C., and McFiggans, G. B.: The effect of photochemical
7 ageing and initial precursor concentration on the composition and hygroscopic properties of
8 beta-caryophyllene secondary organic aerosol, *Atmos. Chem. Phys.*, 12, 6417-6436,
9 10.5194/acp-12-6417-2012, 2012.

10 Bateman, A. P., Nizkorodov, S. A., Laskin, J., and Laskin, A.: Time-resolved molecular
11 characterization of limonene/ozone aerosol using high-resolution electrospray ionization mass
12 spectrometry, *Phys. Chem. Chem. Phys.*, 11, 7931-7942, 10.1039/b905288g, 2009.

13 Bertram, A. K., Martin, S. T., Hanna, S. J., Smith, M. L., Bodsworth, A., Chen, Q., Kuwata, M.,
14 Liu, A., You, Y., and Zorn, S. R.: Predicting the relative humidities of liquid-liquid phase
15 separation, efflorescence, and deliquescence of mixed particles of ammonium sulfate, organic
16 material, and water using the organic-to-sulfate mass ratio of the particle and the oxygen-to-
17 carbon elemental ratio of the organic component, *Atmos. Chem. Phys.*, 11, 10995-11006, Doi
18 10.5194/acp-11-10995-2011, 2011.

19 Bilde, M., and Svenningsson, B.: CCN activation of slightly soluble organics: the importance of
20 small amounts of inorganic salt and particle phase, *Tellus B*, 56, 128-134, Doi 10.1111/j.1600-
21 0889.2004.00090.x, 2004.

22 Brunamonti, S., Krieger, U. K., Marcolli, C., and Peter, T.: Redistribution of black carbon in
23 aerosol particles undergoing liquid-liquid phase separation, *Geophys. Res. Lett.*, 42, 2532-
24 2539, Doi 10.1002/2014gl062908, 2015.

25 Chen, Q., Farmer, D. K., Schneider, J., Zorn, S. R., Heald, C. L., Karl, T. G., Guenther, A., Allan,
26 J. D., Robinson, N., Coe, H., Kimmel, J. R., Pauliquevis, T., Borrmann, S., Poschl, U.,
27 Andreae, M. O., Artaxo, P., Jimenez, J. L., and Martin, S. T.: Mass spectral characterization

1 of submicron biogenic organic particles in the Amazon Basin, *Geophys. Res. Lett.*, 36, Art
2 L2080610.1029/2009gl039880, 2009.

3 Chen, Q., Liu, Y. J., Donahue, N. M., Shilling, J. E., and Martin, S. T.: Particle-Phase Chemistry
4 of Secondary Organic Material: Modeled Compared to Measured O:C and H:C Elemental
5 Ratios Provide Constraints, *Environ. Sci. Technol.*, 45, 4763-4770, 10.1021/es104398s, 2011.

6 Chhabra, P. S., Ng, N. L., Canagaratna, M. R., Corrigan, A. L., Russell, L. M., Worsnop, D. R.,
7 Flagan, R. C., and Seinfeld, J. H.: Elemental composition and oxidation of chamber organic
8 aerosol, *Atmos. Chem. Phys.*, 11, 8827-8845, 10.5194/acp-11-8827-2011, 2011.

9 Ciobanu, V. G., Marcolli, C., Krieger, U. K., Weers, U., and Peter, T.: Liquid-Liquid Phase
10 Separation in Mixed Organic/Inorganic Aerosol Particles, *J. Phys. Chem. A.*, 113, 10966-
11 10978, Doi 10.1021/Jp905054d, 2009.

12 Dusek, U., Frank, G. P., Massling, A., Zeromskiene, K., Iinuma, Y., Schmid, O., Helas, G., Hennig,
13 T., Wiedensohler, A., and Andreae, M. O.: Water uptake by biomass burning aerosol at sub-
14 and supersaturated conditions: closure studies and implications for the role of organics, *Atmos.*
15 *Chem. Phys.*, 11, 9519-9532, 10.5194/acp-11-9519-2011, 2011.

16 Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud droplets
17 and aqueous particles (aqSOA): a review of laboratory, field and model studies, *Atmos. Chem.*
18 *Phys.*, 11, 11069-11102, DOI 10.5194/acp-11-11069-2011, 2011.

19 Frosch, M., Bilde, M., Nenes, A., Praplan, A. P., Juranyi, Z., Dommen, J., Gysel, M., Weingartner,
20 E., and Baltensperger, U.: CCN activity and volatility of beta-caryophyllene secondary
21 organic aerosol, *Atmos. Chem. Phys.*, 13, 2283-2297, 10.5194/acp-13-2283-2013, 2013.

22 Ganbavale, G., Zuend, A., Marcolli, C., and Peter, T.: Improved AIOMFAC model
23 parameterisation of the temperature dependence of activity coefficients for aqueous organic
24 mixtures, *Atmos. Chem. Phys.*, 15, 447-493, 10.5194/acp-15-447-2015, 2015.

25 Gomis-Yagues, V., Ruiz-Bevia, F., Ramos-Nofuentes, M., and Fernandez-Torres, M. J.: The
26 influence of the temperature on the liquid-liquid equilibrium of the ternary system 1-butanol-
27 1-propanol-water, *Fluid Phase Equilibr.*, 149, 139-145, Doi 10.1016/S0378-3812(98)00315-
28 X, 1998.

- 1 Good, N., Topping, D. O., Duplissy, J., Gysel, M., Meyer, N. K., Metzger, A., Turner, S. F.,
2 Baltensperger, U., Ristovski, Z., Weingartner, E., Coe, H., and McFiggans, G.: Widening the
3 gap between measurement and modelling of secondary organic aerosol properties?, *Atmos.*
4 *Chem. Phys.*, 10, 2577-2593, 10.5194/acp-10-2577-2010, 2010.
- 5 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen,
6 J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann,
7 T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W.,
8 McFiggans, G., Mentel, T. F., Monod, A., Prevot, A. S. H., Seinfeld, J. H., Surratt, J. D.,
9 Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic
10 aerosol: current and emerging issues, *Atmos. Chem. Phys.*, 9, 5155-5236, 2009.
- 11 Heald, C. L., Kroll, J. H., Jimenez, J. L., Docherty, K. S., DeCarlo, P. F., Aiken, A. C., Chen, Q.,
12 Martin, S. T., Farmer, D. K., and Artaxo, P.: A simplified description of the evolution of
13 organic aerosol composition in the atmosphere, *Geophys. Res. Lett.*, 37, -, Artn L08803, Doi
14 10.1029/2010gl042737, 2010.
- 15 Heaton, K. J., Dreyfus, M. A., Wang, S., and Johnston, M. V.: Oligomers in the early stage of
16 biogenic secondary organic aerosol formation and growth, *Environ. Sci. Technol.*, 41, 6129-
17 6136, 10.1021/es070314n, 2007.
- 18 Hersey, S. P., Craven, J. S., Metcalf, A. R., Lin, J., Lathem, T., Suski, K. J., Cahill, J. F., Duong,
19 H. T., Sorooshian, A., Jonsson, H. H., Shiraiwa, M., Zuend, A., Nenes, A., Prather, K. A.,
20 Flagan, R. C., and Seinfeld, J. H.: Composition and hygroscopicity of the Los Angeles Aerosol:
21 CalNex, *J. Geophys. Res.-Atmos.*, 118, 3016-3036, 10.1002/jgrd.50307, 2013.
- 22 Hodas, N., Zuend, A., Schilling, K., Berkemeier, T., Shiraiwa, M., Flagan, R. C., and Seinfeld, J.
23 H.: Discontinuities in hygroscopic growth below and above water saturation for laboratory
24 surrogates of oligomers in organic atmospheric aerosols, *Atmos. Chem. Phys.*, 16, 12767-
25 12792, <https://doi.org/10.5194/acp-16-12767-2016>, 2016.
- 26 Hosny, N. A., Fitzgerald, C., Vysniauskas, A., Athanasiadis, A., Berkemeier, T., Uygur, N., Poschl,
27 U., Shiraiwa, M., Kalberer, M., Pope, F. D., and Kuimova, M. K.: Direct imaging of changes

1 in aerosol particle viscosity upon hydration and chemical aging, *Chem. Sci.*, 7, 1357-1367,
2 10.1039/c5sc02959g, 2016.

3 Irwin, M., Good, N., Crosier, J., Choularton, T. W., and McFiggans, G.: Reconciliation of
4 measurements of hygroscopic growth and critical supersaturation of aerosol particles in
5 central Germany, *Atmos. Chem. Phys.*, 10, 11737-11752, 10.5194/acp-10-11737-2010, 2010.

6 Irwin, M., Robinson, N., Allan, J. D., Coe, H., and McFiggans, G.: Size-resolved aerosol water
7 uptake and cloud condensation nuclei measurements as measured above a Southeast Asian
8 rainforest during OP3, *Atmos. Chem. Phys.*, 11, 11157-11174, 10.5194/acp-11-11157-2011,
9 2011.

10 **Jasper, J. J.: The surface tension of pure liquid compounds, *J. Phys. And Chem. Ref. Data*, vol 1,
11 841-1009, Doi: <http://dx.doi.org/10.1063/1.3253106>, 1972.**

12 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H.,
13 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M.,
14 Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A.,
15 Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara,
16 P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J.,
17 Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y.,
18 Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell,
19 L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y.
20 M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M.,
21 Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and
22 Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere, *Science*, 326, 1525-1529,
23 Doi 10.1126/science.1180353, 2009.

24 Juranyi, Z., Gysel, M., Duplissy, J., Weingartner, E., Tritscher, T., Dommen, J., Henning, S., Ziese,
25 M., Kiselev, A., Stratmann, F., George, I., and Baltensperger, U.: Influence of gas-to-particle
26 partitioning on the hygroscopic and droplet activation behaviour of alpha-pinene secondary
27 organic aerosol, *Phys. Chem. Chem. Phys.*, 11, 8091-8097, Doi 10.1039/B904162a, 2009.

- 1 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van
2 Dingenen, R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski,
3 Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K.,
4 Vignati, E., Stephanou, E. G., and Wilson, J.: Organic aerosol and global climate modelling:
5 a review, *Atmos. Chem. Phys.*, 5, 1053-1123, 2005.
- 6 Kang, E., Root, M. J., Toohey, D. W., and Brune, W. H.: Introducing the concept of Potential
7 Aerosol Mass (PAM), *Atmos. Chem. Phys.*, 7, 5727-5744, 2007.
- 8 Knopf, D. A.: Thermodynamic properties and nucleation processes of upper tropospheric and
9 lower stratospheric aerosol particles, Diss. ETH No. 15103, Zurich, Switzerland, 2003.
- 10 Krieger, U. K., Marcolli, C., and Reid, J. P.: Exploring the complexity of aerosol particle properties
11 and processes using single particle techniques, *Chem. Soc. Rev.*, 41, 6631-6662,
12 10.1039/c2cs35082c, 2012.
- 13 Kundu, S., Fisseha, R., Putman, A. L., Rahn, T. A., and Mazzoleni, L. R.: High molecular weight
14 SOA formation during limonene ozonolysis: insights from ultrahigh-resolution FT-ICR mass
15 spectrometry characterization, *Atmos. Chem. Phys.*, 12, 5523-5536, 10.5194/acp-12-5523-
16 2012, 2012.
- 17 Kuwata, M., and Martin, S. T.: Phase of atmospheric secondary organic material affects its
18 reactivity, *P. Natl. Acad. Sci. USA*, 109, 17354-17359, 10.1073/pnas.1209071109, 2012.
- 19 Lambe, A. T., Chhabra, P. S., Onasch, T. B., Brune, W. H., Hunter, J. F., Kroll, J. H., Cummings,
20 M. J., Brogan, J. F., Parmar, Y., Worsnop, D. R., Kolb, C. E., and Davidovits, P.: Effect of
21 oxidant concentration, exposure time, and seed particles on secondary organic aerosol
22 chemical composition and yield, *Atmos. Chem. Phys.*, 15, 3063-3075, 10.5194/acp-15-3063-
23 2015, 2015.
- 24 Li, Y. J., Liu, P. F., Gong, Z. H., Wang, Y., Bateman, A. P., Bergoend, C., Bertram, A. K., and
25 Martin, S. T.: Chemical Reactivity and Liquid/Nonliquid States of Secondary Organic
26 Material, *Environ Sci Technol*, 49, 13264-13274, 10.1021/acs.est.5b03392, 2015.
- 27 Liu, P. F., Abdelmalki, N., Hung, H. M., Wang, Y., Brune, W. H., and Martin, S. T.: Ultraviolet
28 and visible complex refractive indices of secondary organic material produced by

1 photooxidation of the aromatic compounds toluene and m-xylene, *Atmos. Chem. Phys.*, 15,
2 1435-1446, 10.5194/acp-15-1435-2015, 2015.

3 Liu, P. F., Li, Y. J., Wang, Y., Gilles, M. K., Zaveri, R. A., Bertram, A. K., and Martin, S. T.:
4 Lability of secondary organic particulate matter, *P. Natl. Acad. Sci. USA*, 113, 12643-12648,
5 10.1073/pnas.1603138113, 2016.

6 Liu, P. F., Zhang, Y., and Martin, S. T.: Complex Refractive Indices of Thin Films of Secondary
7 Organic Materials by Spectroscopic Ellipsometry from 220 to 1200 nm, *Environ. Sci.*
8 *Technol.*, 47, 13594-13601, Doi 10.1021/Es403411e, 2013.

9 Marcolli, C., and Krieger, U. K.: Phase changes during hygroscopic cycles of mixed
10 organic/inorganic model systems of tropospheric aerosols, *J. Phys. Chem. A.*, 110, 1881-1893,
11 Doi 10.1021/Jp0556759, 2006.

12 Martin, S. T.: Phase transitions of aqueous atmospheric particles, *Chem. Rev.*, 100, 3403-3453,
13 Doi 10.1021/Cr990034t, 2000.

14 Massoli, P., Lambe, A. T., Ahern, A. T., Williams, L. R., Ehn, M., Mikkila, J., Canagaratna, M.
15 R., Brune, W. H., Onasch, T. B., Jayne, J. T., Petaja, T., Kulmala, M., Laaksonen, A., Kolb,
16 C. E., Davidovits, P., and Worsnop, D. R.: Relationship between aerosol oxidation level and
17 hygroscopic properties of laboratory generated secondary organic aerosol (SOA) particles,
18 *Geophys. Res. Lett.*, 37, Artn L2480110.1029/2010gl045258, 2010.

19 O'Brien, R. E., Wang, B. B., Kelly, S. T., Lundt, N., You, Y., Bertram, A. K., Leone, S. R., Laskin,
20 A., and Gilles, M. K.: Liquid-liquid phase separation in aerosol particles: Imaging at the
21 nanometer scale, *Environ. Sci. Technol.*, 49, 4995-5002, 10.1021/acs.est.5b00062, 2015.

22 Pajunoja, A., Lambe, A. T., Hakala, J., Rastak, N., Cummings, M. J., Brogan, J. F., Hao, L. Q.,
23 Paramonov, M., Hong, J., Prisle, N. L., Malila, J., Romakkaniemi, S., Lehtinen, K. E. J.,
24 Laaksonen, A., Kulmala, M., Massoli, P., Onasch, T. B., Donahue, N. M., Riipinen, I.,
25 Davidovits, P., Worsnop, D. R., Petaja, T., and Virtanen, A.: Adsorptive uptake of water by
26 semisolid secondary organic aerosols, *Geophys. Res. Lett.*, 42, 3063-3068, Doi
27 10.1002/2015gl063142, 2015.

- 1 Pandis, S. N., Harley, R. A., Cass, G. R., and Seinfeld, J. H.: Secondary organic aerosol formation
2 and transport, *Atmos. Environ. a-Gen*, 26, 2269-2282, Doi 10.1016/0960-1686(92)90358-R,
3 1992.
- 4 Pankow, J. F.: Gas/particle partitioning of neutral and ionizing compounds to single and multi-
5 phase aerosol particles. 1. Unified modeling framework, *Atmos. Environ.*, 37, 3323-3333, Doi
6 10.1016/S1352-2310(03)00346-7, 2003.
- 7 Pant, A., Parsons, M. T., and Bertram, A. K.: Crystallization of aqueous ammonium sulfate
8 particles internally mixed with soot and kaolinite: Crystallization relative humidities and
9 nucleation rates, *J. Phys. Chem. A.*, 110, 8701-8709, Doi 10.1021/Jp060985s, 2006.
- 10 Parsons, M. T., Mak, J., Lipetz, S. R., and Bertram, A. K.: Deliquescence of malonic, succinic,
11 glutaric, and adipic acid particles, *J. Geophys. Res.-Atmos.*, 109, -, Artn D06212, Doi
12 10.1029/2003jd004075, 2004.
- 13 Petters, M. D., Kreidenweis, S. M., Snider, J. R., Koehler, K. A., Wang, Q., Prenni, A. J., and
14 Demott, P. J.: Cloud droplet activation of polymerized organic aerosol, *Tellus B*, 58, 196-205,
15 DOI 10.1111/j.1600-0889.2006.00181.x, 2006.
- 16 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth
17 and cloud condensation nucleus activity, *Atmos. Chem. Phys.*, 7, 1961-1971, 2007.
- 18 Petters, M. D., Wex, H., Carrico, C. M., Hallbauer, E., Massling, A., McMeeking, G. R., Poulain,
19 L., Wu, Z., Kreidenweis, S. M., and Stratmann, F.: Towards closing the gap between
20 hygroscopic growth and activation for secondary organic aerosol - Part 2: Theoretical
21 approaches, *Atmos. Chem. Phys.*, 9, 3999-4009, 2009.
- 22 Pickering, S.: LXXI. A study of the properties of some strong solutions, *J. Chem. Soc.*, 63, 998-
23 1027, 1893.
- 24 Pöschl, U., Martin, S. T., Sinha, B., Chen, Q., Gunthe, S. S., Huffman, J. A., Borrmann, S., Farmer,
25 D. K., Garland, R. M., Helas, G., Jimenez, J. L., King, S. M., Manzi, A., Mikhailov, E.,
26 Pauliquevis, T., Petters, M. D., Prenni, A. J., Roldin, P., Rose, D., Schneider, J., Su, H., Zorn,
27 S. R., Artaxo, P., and Andreae, M. O.: Rainforest aerosols as biogenic nuclei of clouds and
28 precipitation in the Amazon, *Science*, 329, 1513-1516, Doi 10.1126/science.1191056, 2010.

1 Prenni, A. J., Petters, M. D., Kreidenweis, S. M., DeMott, P. J., and Ziemann, P. J.: Cloud droplet
2 activation of secondary organic aerosol, *J. Geophys. Res.-Atmos.*, 112, Artn D10223, Doi
3 10.1029/2006jd007963, 2007.

4 Prenni, A. J., Petters, M. D., Kreidenweis, S. M., Heald, C. L., Martin, S. T., Artaxo, P., Garland,
5 R. M., Wollny, A. G., and Poschl, U.: Relative roles of biogenic emissions and Saharan dust
6 as ice nuclei in the Amazon basin, *Nat. Geosci.*, 2, 401-404, Doi 10.1038/Ngeo517, 2009.

7 Rastak, N., A. Pajunoja, J. C. Acosta Navarro, D. G. Partridge, J. Ma, M. Song, A. Kirkevåg, Y.
8 Leong, W. W. Hu, N. F. Taylor, D. R. Collins, K. Cerully, A. Bougagioti, R. Krejci, P. Liu,
9 T. Petäjä, C. Percival, A. M. L. Ekman, A. Nenes, S. T. Martin, J. L. Jimenez, D. O. Topping,
10 A. K. Bertram, A. Zuend, A. Virtanen, and I. Riipinen: Microphysical explanation of the RH-
11 dependent water-affinity of biogenic organic aerosol and its importance for climate, *Geophys.*
12 *Res. Lett.*, accepted, 2017.

13 Raymond, T. M., and Pandis, S. N.: Cloud activation of single-component organic aerosol particles,
14 *J. Geophys. Res.-Atmos.*, 107, Artn 478710.1029/2002jd002159, 2002.

15 Reid, J. P., Dennis-Smith, B. J., Kwamena, N. O. A., Miles, R. E. H., Hanford, K. L., and Homer,
16 C. J.: The morphology of aerosol particles consisting of hydrophobic and hydrophilic phases:
17 hydrocarbons, alcohols and fatty acids as the hydrophobic component, *Phys. Chem. Chem.*
18 *Phys.*, 13, 15559-15572, Doi 10.1039/C1cp21510h, 2011.

19 Reinhardt, A., Emmenegger, C., Gerrits, B., Panse, C., Dommen, J., Baltensperger, U., Zenobi, R.,
20 and Kalberer, M.: Ultrahigh mass resolution and accurate mass measurements as a tool to
21 characterize oligomers in secondary organic aerosols, *Anal. Chem.*, 79, 4074-4082,
22 10.1021/ac062425v, 2007.

23 Renbaum-Wolff, L., Song, M. J., Marcolli, C., Zhang, Y., Liu, P. F. F., Grayson, J. W., Geiger, F.
24 M., Martin, S. T., and Bertram, A. K.: Observations and implications of liquid-liquid phase
25 separation at high relative humidities in secondary organic material produced by alpha-pinene
26 ozonolysis without inorganic salts, *Atmos. Chem. Phys.*, 16, 7969-7979, 10.5194/acp-16-
27 7969-2016, 2016.

- 1 Robinson, E. S., Saleh, R., and Donahue, N. M.: Organic aerosol mixing observed by single-
2 particle mass spectrometry, *J. Phys. Chem. A.*, 117, 13935-13945, 10.1021/jp405789t, 2013.
- 3 Schill, G. P., and Tolbert, M. A.: Heterogeneous ice nucleation on phase-separated organic-sulfate
4 particles: effect of liquid vs. glassy coatings, *Atmos. Chem. Phys.*, 13, 4681-4695,
5 10.5194/acp-13-4681-2013, 2013.
- 6 Shilling, J. E., Chen, Q., King, S. M., Rosenoern, T., Kroll, J. H., Worsnop, D. R., DeCarlo, P. F.,
7 Aiken, A. C., Sueper, D., Jimenez, J. L., and Martin, S. T.: Loading-dependent elemental
8 composition of alpha-pinene SOA particles, *Atmos. Chem. Phys.*, 9, 771-782, 2009.
- 9 Shrestha, M., Zhang, Y., Ebben, C. J., Martin, S. T., and Geiger, F. M.: Vibrational sum frequency
10 generation spectroscopy of secondary organic material produced by condensational growth
11 from alpha-pinene ozonolysis, *J. Phys. Chem. A.*, 117, 8427-8436, Doi 10.1021/Jp405065d,
12 2013.
- 13 Smith, M. L., Bertram, A. K., and Martin, S. T.: Deliquescence, efflorescence, and phase
14 miscibility of mixed particles of ammonium sulfate and isoprene-derived secondary organic
15 material, *Atmos. Chem. Phys.*, 12, 9613-9628, 10.5194/acp-12-9613-2012, 2012.
- 16 Song, M., Marcolli, C., Krieger, U. K., Zuend, A., and Peter, T.: Liquid-liquid phase separation
17 and morphology of internally mixed dicarboxylic acids/ammonium sulfate/water particles,
18 *Atmos. Chem. Phys.*, 12, 2691-2712, Doi 10.5194/acp-12-2691-2012, 2012b.
- 19 Song, M., Marcolli, C., Krieger, U. K., Zuend, A., and Peter, T.: Liquid-liquid phase separation in
20 aerosol particles: Dependence on O:C, organic functionalities, and compositional complexity,
21 *Geophys. Res. Lett.*, 39, Artn L19801, Doi 10.1029/2012gl052807, 2012a.
- 22 Song, M. J., Marcolli, C., Krieger, U. K., Lienhard, D. M., and Peter, T.: Morphologies of mixed
23 organic/inorganic/aqueous aerosol droplets, *Faraday Discuss.*, 165, 289-316, Doi
24 10.1039/C3fd00049d, 2013.
- 25 Song, M., Liu, P. F., Hanna, S. J., Li, Y. J., Martin, S. T., and Bertram, A. K.: Relative humidity-
26 dependent viscosities of isoprene-derived secondary organic material and atmospheric
27 implications for isoprene-dominant forests, *Atmos. Chem. Phys.*, 15, 5145-5159, Doi
28 10.5194/acp-15-5145-2015, 2015.

- 1 Spracklen, D. V., Jimenez, J. L., Carslaw, K. S., Worsnop, D. R., Evans, M. J., Mann, G. W.,
2 Zhang, Q., Canagaratna, M. R., Allan, J., Coe, H., McFiggans, G., Rap, A., and Forster, P.:
3 Aerosol mass spectrometer constraint on the global secondary organic aerosol budget, *Atmos.*
4 *Chem. Phys.*, 11, 12109-12136, 10.5194/acp-11-12109-2011, 2011.
- 5 Takahama, S., Schwartz, R. E., Russell, L. M., Macdonald, A. M., Sharma, S., and Leaitch, W. R.:
6 Organic functional groups in aerosol particles from burning and non-burning forest emissions
7 at a high-elevation mountain site, *Atmos. Chem. Phys.*, 11, 6367-6386, Doi 10.5194/acp-11-
8 6367-2011, 2011.
- 9 Tiryaki, A., Guruz, G., and Orbey, H.: Liquid-Liquid Equilibria of Ternary-Systems of Water Plus
10 Acetone and C-5-Alcohol and C-8-Alcohol at 298-K, 303-K and 308-K, *Fluid Phase Equilib.*,
11 94, 267-280, Doi 10.1016/0378-3812(94)87061-6, 1994.
- 12 Veghte, D. P., Altaf, M. B., and Freedman, M. A.: Size Dependence of the Structure of Organic
13 Aerosol, *J. Am. Chem. Soc.*, 135, 16046-16049, 10.1021/ja408903g, 2013.
- 14 Ye, Q., Robinson, E. S., Ding, X., Ye, P. L., Sullivan, R. C., and Donahue, N. M.: Mixing of
15 secondary organic aerosols versus relative humidity, *P. Natl. Acad. Sci. USA*, 113, 12649-
16 12654, 10.1073/pnas.1604536113, 2016.
- 17 You, Y., Renbaum-Wolff, L., and Bertram, A. K.: Liquid-liquid phase separation in particles
18 containing organics mixed with ammonium sulfate, ammonium bisulfate, ammonium nitrate
19 or sodium chloride, *Atmos. Chem. Phys.*, 13, 11723-11734, 10.5194/acp-13-11723-2013,
20 2013.
- 21 You, Y., Smith, M. L., Song, M. J., Martin, S. T., and Bertram, A. K.: Liquid-liquid phase
22 separation in atmospherically relevant particles consisting of organic species and inorganic
23 salts, *Int. Rev. Phys. Chem.*, 33, 43-77, 10.1080/0144235X.2014.890786, 2014.
- 24 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I., Alfarra, M. R.,
25 Takami, A., Middlebrook, A. M., Sun, Y. L., Dzepina, K., Dunlea, E., Docherty, K., DeCarlo,
26 P. F., Salcedo, D., Onasch, T., Jayne, J. T., Miyoshi, T., Shimono, A., Hatakeyama, S.,
27 Takegawa, N., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian,
28 K., Williams, P., Bower, K., Bahreini, R., Cottrell, L., Griffin, R. J., Rautiainen, J., Sun, J. Y.,

1 Zhang, Y. M., and Worsnop, D. R.: Ubiquity and dominance of oxygenated species in organic
2 aerosols in anthropogenically-influenced Northern Hemisphere midlatitudes, *Geophys. Res.*
3 *Lett.*, 34, Artn L13801, Doi 10.1029/2007gl029979, 2007.

4 Zhao, D. F., Buchholz, A., Kortner, B., Schlag, P., Rubach, F., Fuchs, H., Kiendler-Scharr, A.,
5 Tillmann, R., Wahner, A., Watne, A. K., Hallquist, M., Flores, J. M., Rudich, Y., Kristensen,
6 K., Hansen, A. M. K., Glasius, M., Kourtchev, I., Kalberer, M., and Mentel, T. F.: Cloud
7 condensation nuclei activity, droplet growth kinetics, and hygroscopicity of biogenic and
8 anthropogenic secondary organic aerosol (SOA), *Atmos. Chem. Phys.*, 16, 1105-1121,
9 10.5194/acp-16-1105-2016, 2016.

10 Zuend, A., Marcolli, C., Peter, T., and Seinfeld, J. H.: Computation of liquid-liquid equilibria and
11 phase stabilities: implications for RH-dependent gas/particle partitioning of organic-inorganic
12 aerosols, *Atmos. Chem. Phys.*, 10, 7795-7820, Doi 10.5194/acp-10-7795-2010, 2010.

13 Zuend, A., and Seinfeld, J. H.: Modeling the gas-particle partitioning of secondary organic aerosol:
14 the importance of liquid-liquid phase separation, *Atmos. Chem. Phys.*, 12, 3857-3882,
15 10.5194/acp-12-3857-2012, 2012.

16

1 Table 1. Experimental conditions for production and collection of SOM particles by
 2 ozonolysis. Included is the measured separation relative humidity (SRH) upon moistening
 3 and **mixing** relative humidity (MRH) upon drying of the collected particles. Particles were
 4 collected on hydrophobic substrates using an electrostatic precipitator or single stage impactor.
 5 **The uncertainty in the MRH and SRH values are ± 2.0 % RH, due to the uncertainty in the RH**
 6 **measurements.**

7

SOM sample	VOC conc. (ppm)	O ₃ conc. (ppm)	SOM mass conc. ($\mu\text{g m}^{-3}$)*	Flow rate for SOM particle production (L m^{-1})	Collection time (hour)	Collection method	MRH (%)	SRH (%)
β -caryophyllene 1	0.03	30	15-30	7.0	24	Single stage impactor	92.7	94.9
β -caryophyllene 2	0.03	30	15-30	7.0	46	Single stage impactor	95.0	94.4
β -caryophyllene 3	0.7	12	2000-4000	3.5	6	Electrostatic precipitator	90.9	91.5
β -caryophyllene 4	0.7	12	2000-4000	3.5	14	Electrostatic precipitator	93.9	94.1
β -caryophyllene 5	0.7	12	2000-4000	3.5	9	Electrostatic precipitator	93.9	94.1
Limonene 1	0.07	30	80-90	7.0	24	Single stage impactor	95.6	98.7
Limonene 2	0.07	30	80-90	7.0	24	Single stage impactor	97.4	98.8
Limonene 3	2.0	13	7000	3.5	20	Electrostatic precipitator	95.6	95.3
Limonene 4	2.0	13	7000	3.5	20	Electrostatic precipitator	92.7	94.5

8 * Values derived from number-diameter distribution measured by an SMPS and analyzed using a
 9 material density of 1200 kg m^{-3} .

1 Table 2. Experimental conditions for production and collection of SOM produced by photo-
 2 oxidation. Included are the measured separation relative humidity (SRH) upon moistening and
 3 **mixing** relative humidity (MRH) upon drying of the collected particles. Particles were collected
 4 on hydrophobic substrates using an electrostatic precipitator or single stage impactor. SRH = 0 and
 5 MRH = 0 indicates LLPS was not observed during humidity cycles.

SOM sample	VOC conc. (ppm)	O ₃ conc. (ppm)	SOM mass conc. (µg m ⁻³)*	Flow rate for SOM particle production (L m ⁻¹)	Collection time (hour)	Collection method	MRH (%)	SRH (%)
Toluene 1	0.2	30	80-100	7.0	20	Single stage impactor	0	0
Toluene 2	0.2	30	80-100	7.0	24	Single stage impactor	0	0
Toluene 3	1.0	30	600-1000	7.0	48	Electrostatic precipitator	0	0
Toluene 4	1.0	30	600-1000	7.0	48	Electrostatic precipitator	0	0
Toluene 5	1.0	30	600-1000	7.0	96	Electrostatic precipitator	0	0
Toluene 6	1.0	30	600-1000	7.0	96	Electrostatic precipitator	0	0

6 * Values derived from number-diameter distribution measured by an SMPS and analyzed using a
 7 material density of 1200 kg m⁻³.

8

9

1 Table 3. Summary of the LLPS results as well as the average oxygen-to-carbon atomic ratios (O:C)
 2 of the studied SOM particles. The standard deviation (σ) of the separation relative humidity (SRH)
 3 and mixing relative humidity (MRH) is derived from several cycles of RH for different SOM mass
 4 concentrations. SRH = 0 and MRH = 0 indicates phase separation was not observed. The average
 5 O:C values for SOM from toluene photo-oxidation are based on the current study (Section 2.1).
 6 The average O:C values for the other SOM are taken from the literature. Since the average O:C
 7 can depend on oxidant time and oxidation conditions, we chose literature data that were closest to
 8 the experimental conditions used in the LLPS studies (Supplementary Material, Section S1 and
 9 Table S1).

SOM	Average O:C		Average RH (%) $\pm \sigma$	
	Lowest	Highest	MRH	SRH
Ozonolysis of β -caryophyllene	0.36 ^a	0.38 ^a	93.3 \pm 1.7	93.8 \pm 1.3
Ozonolysis of α -pinene	0.42 ^a	0.44 ^a	95.9 \pm 0.8 ^b	96.1 \pm 0.6 ^b
Ozonolysis of limonene	0.34 ^c	0.40 ^c	95.3 \pm 1.9	96.8 \pm 2.2
Photo-oxidation of isoprene	0.52 ^d	0.89 ^a	0 ^e	0 ^e
Photo-oxidation of toluene	1.14 ^f	1.30 ^f	0	0

10 ^a Li et al. (2015); ^b Renbaum-Wolff et al. (2016); ^c Heaton et al. (2007); ^d Lambe et al. (2015); ^e
 11 Rastak et al. (2017); ^f This study.

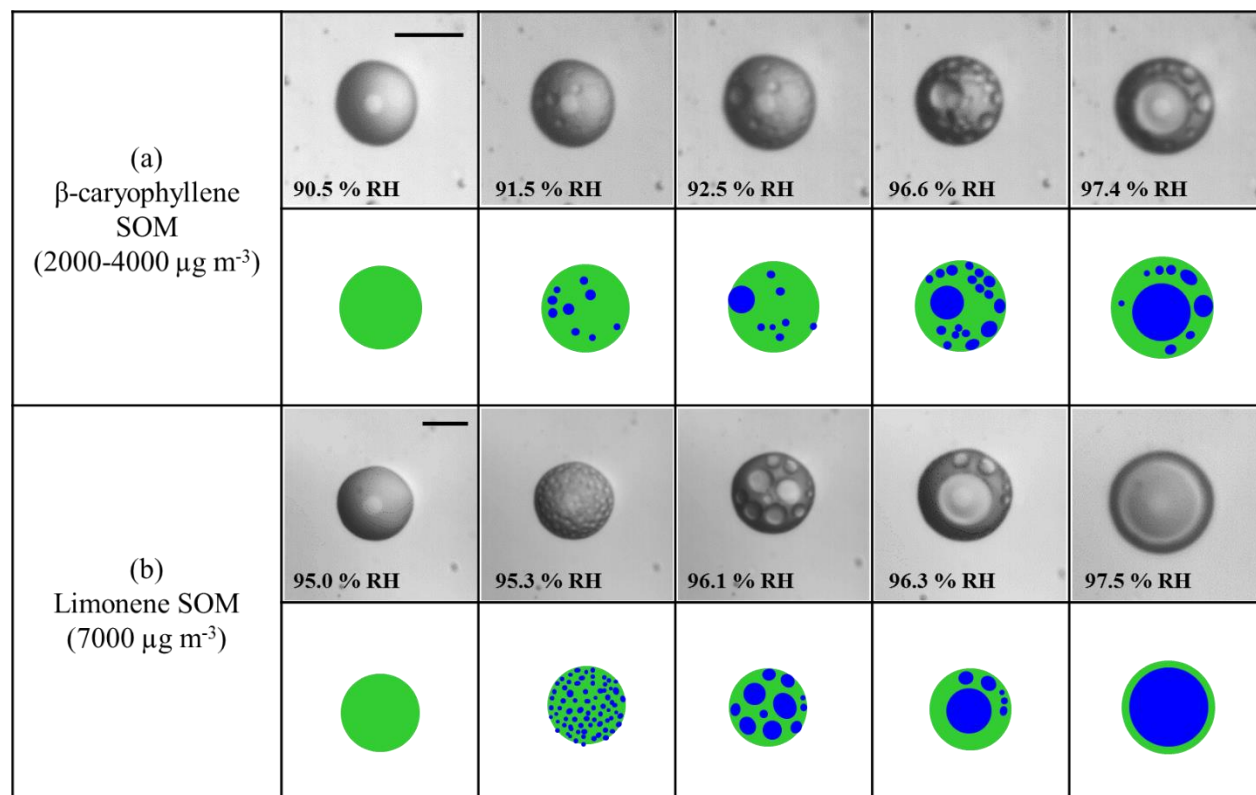
12

1 Table 4. Literature data of measured average O:C, κ_{HGF} , κ_{CCN} , and the difference between κ_{HGF}
 2 and κ_{CCN} , denoted as $\Delta\kappa$, of SOMs. The average O:C values are based on measurements from the
 3 individual studies referenced here. **The experimental conditions for the studies reported in Table**
 4 **4 were not necessarily similar to the experimental conditions for the studies reported in Table 3,**
 5 **even if the same precursor volatile organic compound was used.**

SOM	O:C	κ_{HGF}	κ_{CCN}	$\Delta\kappa$	Reference
		at 90% RH			
Photo-oxidation of α -pinene	0.4	0.04	0.15	0.11	Massoli et al. (2010)
	0.43	0.07	0.16	0.09	Massoli et al. (2010)
	0.45	0.03	0.11	0.08	Pajunoja et al. (2015)
	0.55	0.10	0.12	0.02	Pajunoja et al. (2015)
	0.67	0.14	0.18	0.04	Massoli et al. (2010)
	0.70	0.12	0.13	0.01	Pajunoja et al. (2015)
Photo-oxidation of isoprene	0.86	0.13	0.14	0.01	Pajunoja et al. (2015)
Photo-oxidation of longifolene	0.39	0.02	0.10	0.08	Pajunoja et al. (2015)
	0.56	0.03	0.09	0.06	Pajunoja et al. (2015)
	0.83	0.08	0.10	0.02	Pajunoja et al. (2015)

6

7



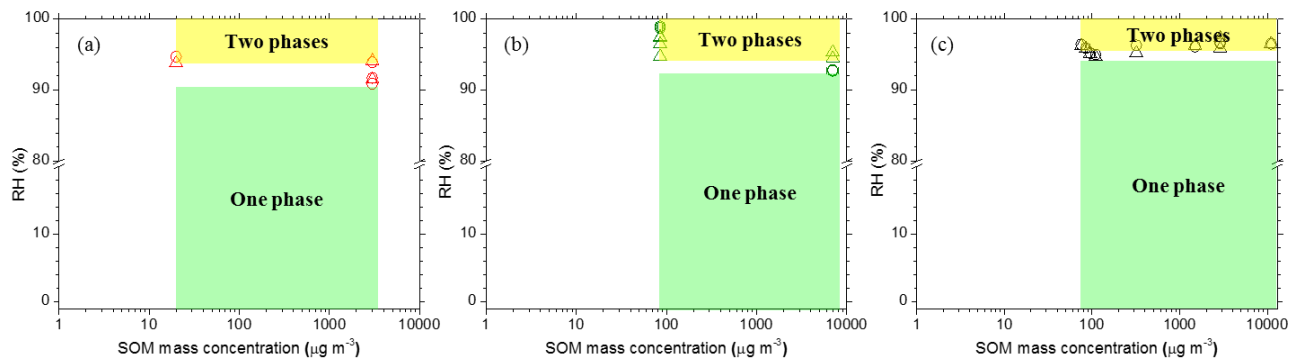
1

2

3 Figure 1. Optical images of SOM particles with increasing RH: (a) β-caryophyllene-derived SOM
 4 for the mass concentrations of 2000 - 4000 μg m⁻³ (β-caryophyllene 3, Table 1) and (b) limonene-
 5 derived SOM for the mass concentrations of 7000 μg m⁻³ (Limonene 3, Table 1). Note that the
 6 light gray circles at the center of the particles are an optical effect due to the hemispherical nature
 7 of the particles. Illustrations are shown below the images for clarity. Green: SOM-rich phase. Blue:
 8 water-rich phase. The scale bar is 20 μm.

9

10



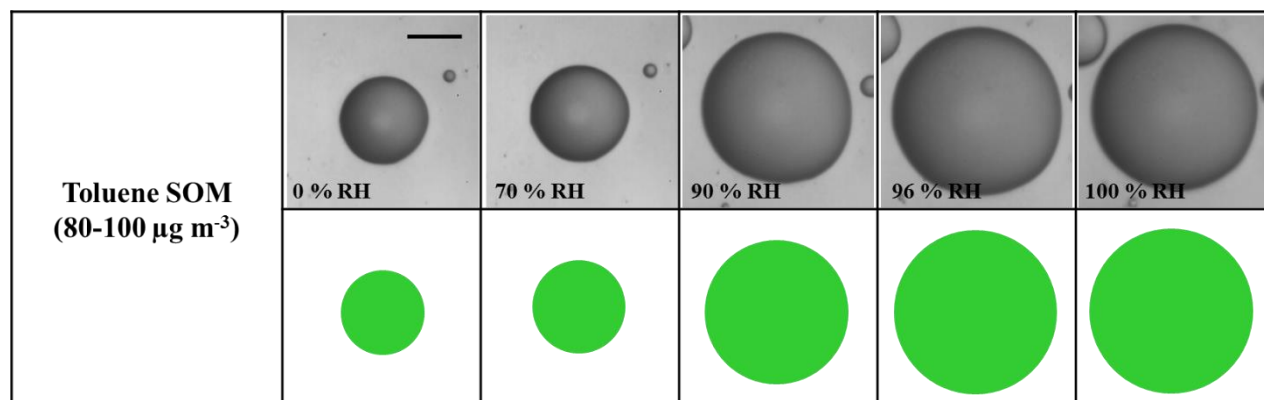
1

2

3 Figure 2. RH at which two phases were observed during humidity cycles of individual particles of
 4 (a) β -caryophyllene-derived SOM and (b) limonene-derived SOM from this study, and (c) α -
 5 pinene-derived SOM from Renbaum-Wolff et al. (2016) as a function of the SOM mass
 6 concentrations. For all panels, circles represent onset of phase separation upon moistening (i.e.
 7 separation relative humidity, SRH) and triangles represent **mixing** of two liquid phases upon
 8 drying (i.e. **mixing** relative humidity, MRH). Yellow shaded region indicates two phases present
 9 and green shaded region indicates one phase prevalent in SOM. **SOM mass concentrations were**
 10 **derived from measured number-diameter distributions and assuming a material density of 1200 kg**
 11 **m^{-3} (Liu et al., 2013).**

12

13



1

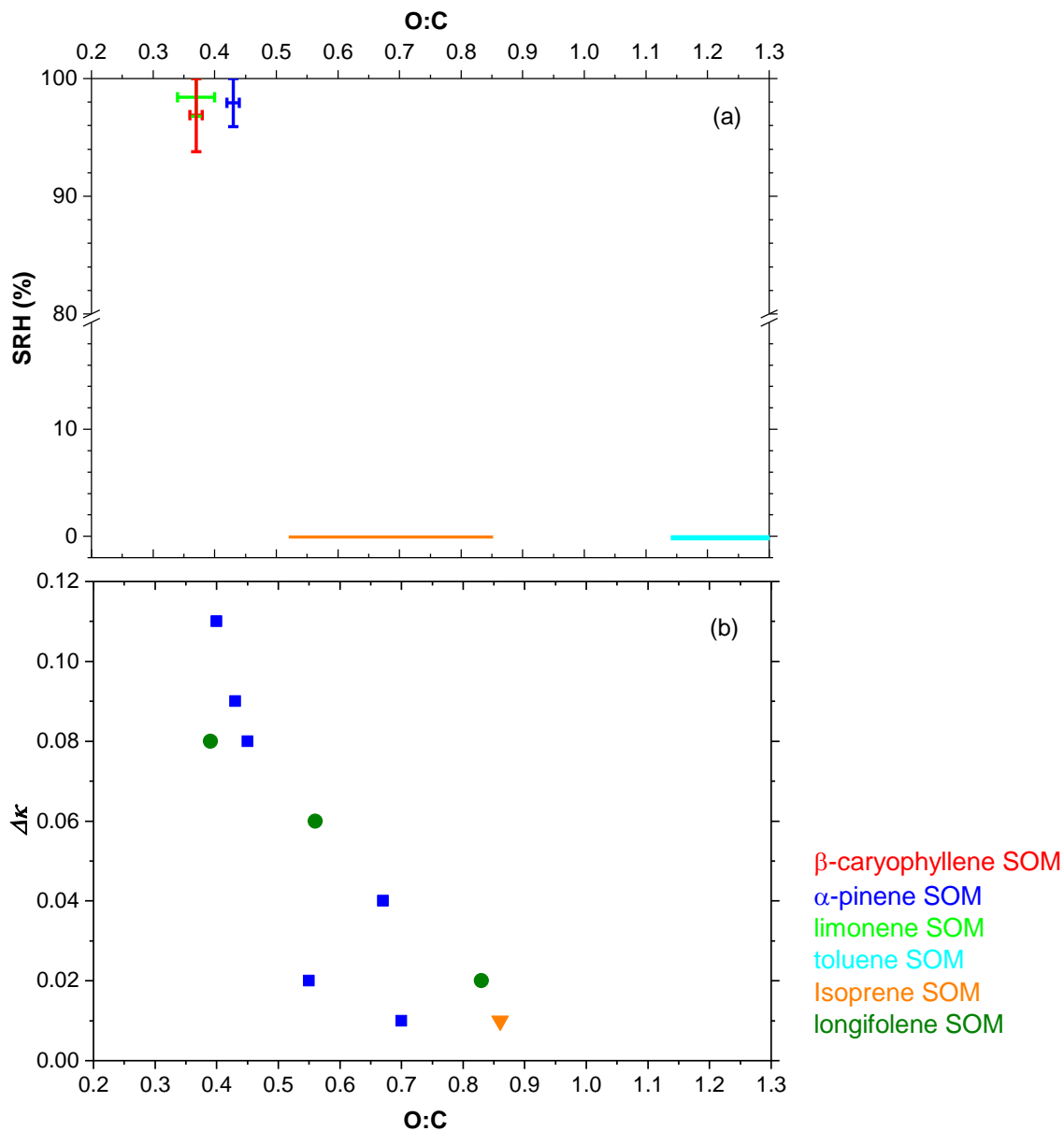
2

3 Figure 3. Optical images of toluene-derived SOM for the particle mass concentrations of 80 - 100
 4 $\mu\text{g m}^{-3}$ (Toluene 2, Table 2) with increasing RH. Illustrations of the images are shown for clarity.

5 Green: SOM-rich phase. Size bar is 20 μm .

6

7



1
2

3 Figure 4. (a) Separation relative humidity (SRH) as a function of the average O:C of the organic
 4 material. Shown are the results from Table 3. β-caryophyllene-derived SOM (red), limonene-
 5 derived SOM (light green), and toluene-derived SOM (cyan) from this study, isoprene-derived
 6 SOM (orange) from Rastak et al. (2017) and α-pinene-derived SOM (blue) from Renbaum-Wolff
 7 et al. (2016) as a function of O:C. RH value of 0 % indicates that LLPS did not occur. The O:C
 8 values of the SOM particles were taken from Table 3. (b) The difference between κ_{HGF} and κ_{CCN} ,

1 denoted as $\Delta\kappa$, as a function of the average O:C of the SOM. Data taken from Table 4. The
2 experimental conditions for the studies reported in Panel (a) were not necessarily similar to the
3 experimental conditions for the studies reported in Panel (b), even if the same precursor volatile
4 organic compound was used.

5