



1 **Aging and hygroscopicity variation of black carbon particles in Beijing measured by a**
2 **quasi-atmospheric aerosol evolution study (QUALITY) chamber**

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13 **Abstract.** Measurements of aging and hygroscopicity variation of black carbon (BC)
14 particles in Beijing were conducted using a 1.2 m³ quasi-atmospheric aerosol evolution study
15 (QUALITY) chamber, which consisted of a bottom flow chamber where ambient air was
16 pulled through continuously and an upper reaction chamber where aging of BC particles
17 occurred. Within the reaction chamber, transmission of the solar ultraviolet irradiation was
18 approximately 50% - 60%, wall loss of primary gaseous pollutants was negligible, and BC
19 exhibited a half-lifetime about 3-7 hours. Typically, equilibrium for the primary gases,
20 temperature, and relative humidity between the reaction chamber and ambient air was
21 established within 1 hour. Rapid growth of BC particles was observed, with an average total
22 growth of 77±33nm and average growth rate of 26±11 nm h⁻¹. Secondary organic aerosol
23 (SOA) accounted for more than 90% of the coating mass. The O/C ratio of SOA was 0.5,
24 lower than the ambient level. The hygroscopic growth factor of BC particles decreased
25 slightly with an initial thin coating layer because of BC reconstruction, but subsequently
26 increased to 1.06-1.08 upon further aging. The κ (kappa) values for BC particles and coating
27 materials were calculated as 0.035 and 0.040 at the subsaturation and supersaturation
28 conditions, respectively, indicating low hygroscopicity of coated SOA on BC particles. Hence,
29 our results indicate that initial photochemical aging of BC particles does not appreciably alter
30 the particle hygroscopicity in Beijing.

31



32 1 Introduction

33 Atmospheric aerosols undergo continuous and complicated transformation during their
34 residence time in the atmosphere. The aging of aerosols is likely resulted from both physical
35 (i.e., coagulation, condensation, equilibrium partitioning, and evaporation) and chemical (i.e.,
36 photochemical gas-phase oxidation and multi-phase reactions) processes (Zhao et al., 2006;
37 Qiu et al., 2013; Zhang et al., 2015). Also, there are typically large variations in the particle
38 properties (i.e., size, mass, chemical composition, morphology, and optical and hygroscopic
39 parameters) during aging, significantly influencing the aerosol impacts on visibility, human
40 health, weather, and climate (Jacobson, 2001; Guo et al., 2014). A better understanding of the
41 aging process of aerosols in the atmosphere is critical in atmospheric and climate research.

42 For example, the scientific interest in the climate effects of black carbon (BC) has
43 remained, since BC is the strongest absorber of visible solar radiation (Wang et al., 2013). BC
44 solar absorption represents a central issue in climate change research, since the synthesis of
45 satellite, in situ, and ground observations shows that the global solar absorption (direct
46 radiative forcing or DRF) by BC is as much as $0.9 \text{ W}\cdot\text{m}^{-2}$, second only to that of CO_2
47 (Jacobson, 2001; Bond et al., 2013; IPCC, 2013). BC also represents an important component
48 of air pollution for large parts of the world (Zhang et al., 2015). The properties of BC are
49 considerably modified during aging, including the size, mass, morphology, and optical and
50 hygroscopic parameters (Khalizov et al., 2009; Xue et al., 2009a). Enhanced light absorption
51 of BC particles during aging not only contributes to atmospheric stabilization and
52 exacerbation of haze formation, but also imposes large positive radiative forcing on climate
53 (Peng et al., 2016). Furthermore, the variation in hygroscopicity during aging also regulates
54 the lifetimes of BC particles. Hygroscopic particles serve efficiently as cloud condensed
55 nuclei (CCN), affecting the formation, longevity and albedo of clouds (Yuan et al., 2008;
56 Wang et al., 2011). Also, deposition of BC particles, via in-cloud scavenging and wet
57 deposition, depends highly on the particle hygroscopicity (Bond et al., 2013). In addition, the
58 hygroscopicity also affects the aqueous-phase reactions of atmospheric pollutants (Ervens et
59 al., 2011; Wang et al., 2016). Previous studies using hygroscopic tandem differential mobility
60 analyzer (H-TDMA) instruments have shown that coating of hydrophilic materials
61 significantly increases the hygroscopic growth factor of BC particles (Saathoff et al., 2003;
62 Khalizov et al., 2009; Guo et al., 2016). The ability of BC particles to form CCN also
63 enhanced after coating of hydrophilic materials (Kuwata et al., 2007; Tritscher et al., 2011;
64 Ma et al., 2013; Wittbom et al., 2014). The activation supersaturation depends on the particle
65 size, hygroscopicity of coating materials, and the coating thickness (Ma et al., 2013). The
66 coating materials in the previous experiments include sulfuric acid (Zhang and Zhang, 2005;
67 Khalizov et al., 2009), oxidation products from biogenic and anthropogenic hydrocarbon
68 species (Saathoff et al., 2003; Ma et al., 2013; Khalizov et al., 2013), or SOA from single
69 emission source (Tritscher et al., 2011). However, there still exist uncertainties for
70 parametrization of the BC lifetime in atmospheric models, because of insufficient constraints
71 on the hydrophobic to hydrophilic conversion of BC particles under variable ambient
72 conditions.

73 Atmospheric field measurements have been performed to evaluate aging of particles on
74 different platforms, e.g., ground, aircraft, and cruise (Moffet and Prather, 2009; DeCarlo et al.,



75 2010; Peng et al., 2014; Liu et al., 2015) and over different spatial scales (intensive
76 campaigns or long-term measurements). Typically, a wide variety of state-of art instruments
77 are employed to characterize the changes of the chemical and physical properties of aerosols.
78 On the other hand, field measurements at fixed sites are affected by transport, local emissions,
79 and chemistry, and quantification of the particle parameters during aging involves complex
80 decoupling of the various processes (Peng et al., 2016). In particular, it is challenging to
81 isolate the chemical processes from those related to meteorology (i.e., transport and mixing)
82 and emissions.

83 The methods of environmental chambers or reactors have been widely employed in
84 atmospheric chemistry research, including photochemical oxidation of volatile organic
85 compounds (VOCs) (Zhang et al., 2000), formation and growth of aerosols (Claeys, 2004;
86 Kalberer, 2004), nucleation of nanoparticles (Zhang et al., 2004; Wang et al., 2010; Zhang et
87 al., 2012), aging of BC particles (Zhang et al., 2008), and cloud formation (Ruehl et al.,
88 2016). Dependent of the scientific objectives, the designs of environmental chambers and
89 reactors vary considerably (Zhang et al., 2015). However, few of the previous experimental
90 methods are able to characterize the evolution of aerosols under the ambient conditions.

91 In this study, we present measurements of aging and hygroscopicity variation of BC
92 particles in Beijing using a quasi-atmospheric aerosol evolution study (QUALITY) chamber
93 (Reed, 2010; Peng et al., 2016). The performance of the QUALITY chamber for mimicking
94 the ambient gaseous concentrations (i.e., the wall loss, and gas mixing rate), ultraviolet
95 transmission, and meteorology parameters (i.e., temperature and relative humidity, RH) has
96 been evaluated.

97 **2 Experimental method**

98 The 1.2 m³ QUALITY chamber was employed to study BC aging under ambient
99 conditions (Fig. 1). The two-layer chamber was comprised of an inner layer of 0.13 mm
100 polytetrafluoro ethylene (PFA) Teflon and an outer rigid 5.6 mm thick acrylic shell (Cryo
101 Industries Acrylite, OP-4). Both Acrylite OP-4 Acrylic and PFA Teflon allowed for efficient
102 ultraviolet (UV) transmission in UV-B (280-315 nm) and UV-A (315-400 nm) ranges. When
103 exposed to sunlight, the UV light transmitted through the chamber wall and initiated
104 photochemical reactions inside the chamber.

105 The two individual subdivisions of the QUALITY chamber included a bottom flow
106 chamber, where ambient air was pulled through continuously over each experiment, and an
107 upper reaction chamber, where aging of BC particles occurred (Fig. 1). The two chambers
108 were separated by a 5 μm thick semi-permeable expanded polytetrafluoroethylene (e-PTFE)
109 membrane with high degrees of chemical resistivity, microporosity, nonpolarity, and thermal
110 stability. Gaseous species penetrated the membrane by either bulk gas flow or diffusion. The
111 permeability of the e-PTFE was greater than 90 % for nearly all the tested volatile organic
112 components (Fig. S1) and other gaseous pollutants, i.e., SO₂, NO_x, O₃ and CO. Particles, on
113 the other hand, were blocked from penetration into the reaction chamber. The filtration
114 efficiency of the e-PTFE membrane was measured to be over 99.5% for particles larger than
115 15 nm. During each chamber experiment, ambient air was pulled through the flow chamber
116 continuously and gases in lower chamber permeated through the membrane into the reaction



117 chamber. Hence, an environment that continuously captured the ambient gas concentrations
118 without the presence of ambient particles was created inside the reaction chamber. Since the
119 chamber was continuously exposed to ambient gas concentrations during experiments, gases
120 lost due to reaction, deposition or adsorption to the seed aerosols within the reaction chamber
121 were steadily replenished by the exchange with the flow chamber. Several sampling ports
122 were set at the side of the reaction chamber for injection of seed particles or sampling during
123 experiment.

124 Seed particles were introduced in the reaction chamber via an injection line (Fig. S2). To
125 investigate the growth of BC particles under ambient condition, monodisperse BC particles
126 were injected into the QUALITY chamber. BC particles were generated by incomplete
127 combustion of propane fuel in a custom-made laminar diffusion burner (Santoro et al., 1983;
128 Qiu et al., 2012). The aerosol stream sequentially passed through a 300°C heater to evaporate
129 the semi-volatile organic compounds in the particle phase, a Nafion dryer to remove excess
130 moisture in the flow, and four one-meter-long cylinder containing both alumina spherules
131 coated with potassium permanganate and activated carbon to remove all the gaseous
132 pollutants (i.e., VOCs, H₂S, SO₂, NO_x, O₃). The measured removal efficiency in the cylinders
133 for SO₂, NO and NO₂ were 99.2%, 100%, and 99.9%, respectively. The aerosol stream was
134 then introduced through an ionizer and into a differential mobility analyzer (DMA, model
135 3081, TSI, Inc.) with stable voltage to create a monodisperse BC particles flow.

136 A suite of high time resolution state-of-the-art aerosol instruments simultaneously
137 measured a comprehensive set of BC properties throughout the BC aging process (Table S1).
138 The particle diameter, mass, chemical composition, hygroscopicity and ability as cloud
139 condensation nuclei (CCN) were measurement by a scanning mobility particle sizer (SMPS),
140 a differential mobility analyzer–aerosol particle mass analyzer (DMA-APM) system, a
141 high-resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS), a humidified
142 tandem differential mobility analyzer (HTDMA) system, and a cloud condensation nuclei
143 counter (CCNC), respectively. Detail information of the instruments is provided in the
144 supplementary material and previous publications (DeCarlo et al., 2006; Khalizov et al.,
145 2009). Specific measurement procedures in this study are discussed below.

146 **DMA-APM measurement.** The DMA-APM was used to measure the effective density of
147 BC particles (Pagels et al., 2009). Before any DMA-APM measurement, a SMPS scan was
148 made to obtain the size distribution of particles inside the chamber. The particle size
149 distribution was then fitted with a lognormal Gaussian distribution to derive the peak
150 diameter. During a DMA-APM measurement, the aerosol flow passed through DMA with a
151 fixed voltage to select particles with a fixed diameter. The APM then measured the mass
152 distribution of the selected particles with the same diameter, and the effective density of these
153 particles was obtained by fitting the mass distribution with a normal Gaussian distribution.

154 **DMA-CCN measurement.** Similar to the density measurements, CCN activation was also
155 measured on the basis of monodisperse particles with a peak diameter. A DMA with a fixed
156 voltage selected particles with a peak diameter. Both a CPC and a CCN counter were placed
157 in parallel after the DMA to simultaneously measure the total particle number concentration
158 (N_{cn}) as well as the activated number concentration (N_{ccn}) at a fixed supersaturation. The



159 activation fraction of the BC particles with peak diameter is calculated as:

$$160 \quad f_{CCN} = \frac{N_{CN}}{N_{CCN}} \quad (1)$$

161 Several gradients of supersaturation were set for the CCN counter, with each one being 6
162 min. This method yielded a steeper curve of the CCN activation rate, which was employed to
163 estimate the particle diameter with 50% activation fraction (D_{50}) and the kappa value.

164 Prior to each experiment, the QUALITY chamber was flushed by zero air for more than
165 40 hours to ensure a clean condition and covered with two layers of anti-UV cloth to shield it
166 from sunlight. In the beginning of each experiment, monodisperse BC particles were
167 introduced into the chamber. The injection of BC particles typically lasted for 1 to 2 hours.
168 During the injection period, zero air passed through the bottom chamber continually to
169 remove any possible remaining gaseous pollutant that were removed by the activated carbon.
170 After the injection, ambient air was pulled through the bottom chamber at a flow rate of about
171 50 L min⁻¹ for at least half an hour in order to produce a quasi-ambient condition inside the
172 chamber. Finally, the anti-UV cloth was removed, and BC particles underwent aging inside
173 the reaction chamber. A charged zero air stream continuously passed through the space
174 between the two chamber layers to reduce particle wall loss. Various properties of BC
175 particles, including the particle diameter, mass, chemical composition, hygroscopicity, and
176 optical coefficients, were simultaneously measured by a suite of state-of-art aerosol
177 instruments in every 0.5 - 1 hour. Ambient particles and chamber particles were measured
178 alternately every 30 min. The aging experiments lasted for about 2-6 hours depending on the
179 initial BC concentrations and ambient conditions.

180 The BC aging experiments were conducted from August 18th to October 17th, 2013 at an
181 urban site (PKUERS) located on the campus of Peking University in the northwestern Beijing
182 (39.99°N, 116.31°E) (Hu et al., 2012).

183 **3. Characterization and Validation of the QUALITY chamber**

184 **3.1 Wall loss of gases and aerosols**

185 To evaluate the wall loss of both particles and gases in the QUALITY chamber, different
186 gaseous pollutants and particles were introduced into the chamber separately and the decay of
187 their concentration inside the chamber was measured by gas analyzers and SMPS,
188 respectively. All of the ports connected to the ambient air were closed to ensure an enclosed
189 system in the reaction chamber.

190 Particles with different chemical composition exhibited different wall loss rates.
191 Monodispersed BC particles with different diameters showed a small half-lifetime (τ) of
192 about 4-7 hours (Reed, 2010). Aerosol nucleation also occurred inside the chamber likely
193 from organic species (Zhao et al., 2009), which corresponded to a half time of about 3.5
194 hours (Fig. S3), because the nucleated particles inside the chamber was neutral with a slow
195 electrostatic loss to the wall.

196 Gas species had a longer residence time inside the chamber. Toluene and isoprene did
197 not show obvious wall loss during a two-day experiment (Reed, 2010). O₃, SO₂ and NO_x



198 decreased by 50% inside the chamber after more than 20 hours, suggesting slow loss. Since
199 the loss rate of these primary gaseous pollutants was much slower than the gas exchange rate
200 between the chamber, the loss of gases was replenished by the exchange with the flow
201 chamber.

202 3.2 UV transmission

203 The QUALITY chamber contained two layers of walls, an acrylic shell layer and a PFA
204 Teflon layer. Since the QUALITY chamber used sunlight as the photochemical origin, the
205 transmission spectra of the two-layer walls was of great important for the photochemical
206 reactions inside the reaction chamber.

207 The transmission efficiencies of each material were measured using a Fourier transform
208 infrared spectroscopy (FTIR) system. As shown in Figure 2, the Teflon film exhibited stable
209 transmission efficiency of about 60% in the focused wavelength range. The Acrylic shell,
210 however, showed very low transmission efficiency when the wavelength range was shorter
211 than 270 nm, and high transmission efficiency (nearly 90%) when the wavelength range was
212 longer than 300 nm. In general, approximately 60% of the UVA irradiation (315-400 nm
213 range) and 50% of the UVB irradiation (280-315 nm range) penetrated through the chamber
214 walls, allowing photochemical processes to take place in the upper reaction chamber.

215 A NO₂-photolysis experiment was also conducted to characterize the UV transmission of
216 QUALITY chamber. NO₂ was introduced inside the chamber at a clean and sunny day.
217 Reactions among NO₂, NO and O₃ will occur as the following equations:



221 By simultaneously monitoring the concentrations of NO₂, NO and O₃ inside the chamber,
222 the photolysis constant of NO₂, J(NO₂), inside the chamber was estimated. The photolysis
223 constant of NO₂ inside the chamber was on average 55% of that in the ambient air, in
224 agreement with the transmission spectra measurement, further confirming that the two-layer
225 chamber walls allowed 50% - 60% of solar irradiation in the UV range into the reaction
226 chamber.

227 3.3 Exchange of gases between the reaction chamber and ambient air

228 Additional experiments were conducted to characterize the exchange time scale for gases
229 in the QUALITY chamber. The chamber was firstly cleaned and flushed with zero air for 40
230 hours. Ambient air was then pulled through the bottom chamber (Fig. 3), and the
231 concentrations of gaseous pollutants, including O₃, NO_x, CO, and SO₂, were measured
232 alternatively in the ambient air and in the upper chamber. In the beginning of the experiment
233 the concentrations of all gaseous species were lower than those in the ambient air. After
234 ambient air was pulled through the bottom chamber (labeled as the black dots in Fig. 3), the
235 concentrations of the gaseous pollutants in the reaction chamber increased sharply. For



236 example, the CO concentration inside the chamber was approximate 70% of that the ambient
237 concentration after 30-min mixing. The concentrations inside the chamber and ambient air
238 exhibited little difference after 1-hour mixing (Fig. 3c). Hence, the QUALITY chamber well
239 replicated the ambient gas concentrations. The gas exchange rate between the bottom and
240 reaction chambers was calculated to be approximately 0.06 min^{-1} .

241 3.4 Temperature and RH

242 The greenhouse effect for an outdoor chamber typically increases the temperature and
243 decreases the RH inside the chamber. For the QUALITY chamber, however, heat produced
244 by the greenhouse effect inside the chamber was effectively taken away as the ambient air
245 continuously passed through the bottom chamber and exchanged with air in the upper
246 chamber. As illustrated in Figure 4, there was little difference in temperature or RH inside
247 and outside the chamber, when the chamber experiments lasted for more than 1 hour,
248 suggesting that the QUALITY chamber effectively captured the ambient temperature and
249 RH.

250 3.5 Sulfuric acid production

251 Though gaseous pollutants such as SO_2 , O_3 , NO_x and VOCs penetrated into the reaction
252 chamber from the bottom chamber through the semi-permeable membrane, low volatile and
253 sticky gases, i.e., sulfuric acid, were unlikely to penetrate through the membrane (Fortner et
254 al., 2004). To characterize sulfuric acid production inside the chamber, we conducted a
255 special experiment by pulling ambient air into the bottom chamber, while injecting SO_2
256 directly into the reaction chamber. The experiment was conducted around noon on a clean
257 day when the O_3 concentration was around 50 ppb. A custom-made atmospheric pressure-ion
258 drift chemical ionization mass spectrometry (AP-ID-CIMS) (Fortner et al., 2004; Zheng et al.,
259 2010) was used to directly measure the concentration of gaseous sulfuric acid. Fig. 5 shows a
260 good correlation between SO_2 and gaseous H_2SO_4 inside the chamber, suggesting that the
261 QUALITY chamber well simulated the formation of low volatile gaseous species and hence
262 the photochemical processes.

263 4. BC evolution in the QUALITY chamber

264 Time series of the ambient PM, gas, and meteorology parameters of ambient air during
265 chamber experiment period are illustrated in Figure 6. Except for the last experiment
266 conducted on October 17th, all experiments were conducted between August 18th, 2013 and
267 September 21st, 2013. During this period in Beijing, strong photochemical reactions and
268 frequent heavy pollution events occurred (Huang et al., 2010; Zhao et al., 2013). Temperature
269 and RH during this period (August 18th to September 21st) varied from 24°C to 38°C and
270 from 20 % to 90 %, respectively. The average concentration of $\text{PM}_{2.5}$, SO_2 and NO_x were 60
271 $\pm 45 \mu\text{g m}^{-3}$, $3.2 \pm 2.6 \text{ ppb}$ and $33.9 \pm 20.9 \text{ ppb}$, respectively.

272 The red shaded areas in Figure 6 represent the period of the nine chamber experiments.
273 In general, chamber experiments were conducted in the afternoon of relatively clean and
274 sunny days, when strong solar radiation led to fast photochemical reactions. There were two
275 experiments conducted under polluted days, e.g., the experiments on October 22nd and
276 September 11th. Table 1 summaries the conditions of the experiments. Totally, 10 BC aging



277 experiments were conducted, including four experiments using BC particles with initial
 278 mobility diameter of 100 nm, three experiments using 150 nm particles, and three using 220
 279 nm BC particles. The average concentrations of PM_{2.5} and NO_x over each chamber
 280 experiment were only 9 to 69 μg m⁻³ and 9 to 41 ppb, respectively. The concentrations of
 281 VOCs, such as toluene and m/p-xylene were relatively low during most of the experiments
 282 compared with severe pollution episodes in Beijing (Guo et al., 2014). J(O¹D) and O₃
 283 exhibited higher values during the chamber experiments. The average J(O¹D) values and O₃
 284 concentrations ranged from 3.2 × 10⁻⁶ s⁻¹ to 21.1 × 10⁻⁶ s⁻¹ and 26 to 92 ppb, respectively.

285 4.1 BC growth

286 To quantify the growth of BC particles, several parameters were used to describe the
 287 properties of BC particle, including the effective density, mobility diameter and mass
 288 equivalent diameter. Material density (ρ_m) is the average density of the solid and liquid
 289 material in the particle. Assuming that the volume of the species coexisting in an individual
 290 particle does not change upon mixing, the density of an internally mixed particle is calculated
 291 using the material densities and mass concentrations of particulate constituents (DeCarlo et
 292 al., 2004; Pagels et al., 2009),

$$293 \quad \rho_m = \frac{\sum_{species} m_i}{\sum_{species} V_i} = \frac{\sum_{species} m_i}{\sum_{species} \frac{m_i}{\rho_i}} = \frac{\sum_{species} MC_i}{\sum_{species} \frac{MC_i}{\rho_i}} \quad (5)$$

294 where ρ_i is the material density of species i , V_i is its volume, m_i is its mass, and MC_i is its
 295 mass concentration. This approach is based on the assumption that there is no void space
 296 enclosed within the particle envelope. Hence, the material density is larger than the true
 297 particle density with internal voids in particles. In this study, ρ_m is calculated from the
 298 chemical composition of coating materials measured by AMS. A value of 1.35 for the
 299 material density of SOA formed during chamber experiment was obtained by directly
 300 measuring the density of newly form particles inside the chamber via DMA-APM system.

301 Effective density (ρ_{eff}) is defined as the ratio of the measured particle mass (m_p) to the
 302 particle volume calculated assuming a spherical particle with a diameter equal to the
 303 measured mobility diameter (D_m) (DeCarlo et al., 2004; Pagels et al., 2009; Xue et al.,
 304 2009b):

$$305 \quad \rho_{eff} = \frac{6m_p}{\pi D_m^3} \quad (6)$$

306 In this study, m_p of BC particles was measured by the APM and mobility diameter (D_m)
 307 was measured by the DMA. The effective density reflects the information on both particle
 308 density and shape. If particles are spherical in the absence of internal void, the effective
 309 density equals the material density. If particles are non-spherical, the calculated volume and
 310 volume concentration are larger than the true values, and the effective density is less than true
 311 particle and the material density.

312 Figure 7 exhibits the change of particle density and diameter in three typical BC aging
 313 experiments using BC particles with the initial diameter of 100 nm, 150 nm, and 220 nm



314 (experiment #4, #5, and # 9 in Table 1, respectively). The average $PM_{2.5}$ concentrations in
315 these experiments were 40, 27, and 12 $\mu\text{g m}^{-3}$, respectively, suggesting relatively clean
316 conditions during the experiments. In all three experiments, aging of BC particles occurred
317 between 13:00-14:00 and 17:00 in the afternoon. The highest $J(\text{O}^1\text{D})$ value varied from 1.7 to
318 $2.4 \times 10^{-5} \text{ s}^{-1}$ and decreased generally over the experiment period. Average O_3 concentrations
319 during the three experiments were 68, 83 and 54 ppb, respectively, indicating strong
320 oxidation during the experiment periods.

321 The initial ρ_{eff} values of BC particles were 0.46, 0.34 and 0.25 g cm^{-3} for particles with
322 the initial diameter of 100 nm, 150 nm, and 220 nm, respectively, indicating fractal BC
323 aggregates with chain-like branches (Zhang et al., 2008). After aging of 1 hour, ρ_{eff} of BC
324 particles in the three experiments increased to approximately 1.2 g cm^{-3} , suggesting that
325 formation of the secondary components changed the morphology from chain-like BC
326 particles into a more compact shape (Zhang et al., 2008; Peng et al., 2016). The morphology
327 change was further confirmed by a decrease of D_m , particularly for larger BC particles that
328 were more fractal than smaller BC particles.

329 As the D_m is largely influenced by the particle morphology, we utilized the parameter of
330 mass equivalent diameter (D_{me}) to describe the growth of particles. Based on the mobility
331 diameter (D_m), material density (ρ_m) and effective density (ρ_{eff}), the D_{me} is calculated
332 assuming that particles are compact and with a spherical morphology (DeCarlo et al., 2004):

$$333 \quad D_{me} = \sqrt[3]{\frac{\rho_{eff}}{\rho_m}} D_m \quad (7)$$

334 The change in the mass equivalent diameter (ΔD_{me}) during BC aging is defined as the total
335 coating thickness, and the ratio of the total coating thickness to the initial mass equivalent
336 diameter ($\Delta D_{me}/D_{me,0}$) is defined as the coating fraction.

337 The initial D_{me} of fresh BC particles with initial D_m of 100, 150 and 220 nm were 61, 84
338 and 114 nm, respectively. In contrast to the mobility diameter, D_{me} increased continuously
339 during the entire experiment. After 3-4 hours, D_{me} in the three experiments increased to 133,
340 169 and 197 nm, respectively (Fig. 7), with the average growth rates of 19, 29 and 31 nm h^{-1} .
341 Higher growth rates in D_{me} occurred around noontime, when the $J(\text{O}^1\text{D})$ value was higher and
342 the photochemical reaction was stronger. On the other hand, much less growth rate was
343 observed during late afternoon or with cloud coverage (As shown in Fig. 7c at 15:00),
344 indicating that the growth was driven by photochemical reactions.

345 The increases of the particle density and diameter in all the experiments are summarized
346 in Table 2. Fast aging of BC particles occurred in all experiments. The total growth of D_{me}
347 ranged from 40 nm to 152 nm within 3-6 hours, with an average growth of 73 nm. The
348 average growth rate was $26 \pm 11 \text{ nm h}^{-1}$, demonstrating large secondary production under the
349 ambient conditions in Beijing. The largest growth rate ($\Delta D_{me} = 152 \text{ nm}$) was observed in
350 experiment #8, when solar irradiation was the strongest among all experiments (Table 2).

351 Correlation analysis was made between the average growth rate of BC particles ($\Delta D_{me}/\Delta t$)
352 with O_3 , $PM_{2.5}$, $J(\text{O}^1\text{D})$, and temperature during the different experiments (Fig. S4). The



353 growth rate of BC particles exhibits no correlation with O_3 concentration ($R^2=0.00$), weak
354 negative correlation with $PM_{2.5}$ concentration ($R^2 = 0.25$), and strong positive correlations
355 with $J(O^1D)$ ($R^2 = 0.80$) and temperature ($R^2 = 0.67$), indicating the importance of
356 photochemical production on the BC coating materials.

357 4.2 Chemical composition of coating materials

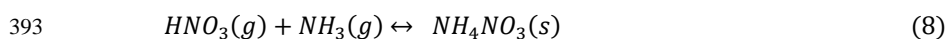
358 Particle composition measurements by AMS during chamber experiments reveal a
359 majority of coating materials (above 90%) as SOA (Fig. 7). The concentration of SOA inside
360 the chamber reached up to $9 \mu g m^{-3}$ in several experiments, suggesting fast formation of SOA
361 via gas phase oxidation of VOCs. The SOA formation in Beijing is likely attributed to a large
362 amount of anthropogenic aromatic VOCs (Peng et al., 2017).

363 The elemental compositions of OA inside the chamber, i.e., the oxygen to carbon (O/C)
364 ratio and the hydrogen to carbon (H/C) ratio, were calculated based on the updated ambient
365 calibrations (He et al., 2011; Canagaratna et al., 2015). The H/C and O/C ratios of organics
366 for coating on BC particles exhibit notable trends during the aging process. Figure 8A shows
367 an example of the evolution of H/C and O/C ratios in experiment #8. The data were corrected
368 for the CO_2 concentration in the chamber, which were introduced into the chamber with BC
369 particles and influenced the abundance of $m/z = 28$ and 44 in the AMS mass spectra. The H/C
370 ratio decreased from 1.73 to 1.45 over six hours. Accordingly, the O/C ratio increased from
371 0.32 to 0.50 during the same time, revealing that further oxidation of SOA occurred in the
372 latter part of the experiment. The lower final O/C ratio in the chamber experiment (0.5) than
373 that under the ambient conditions (Hu et al., 2016) implies that there is oxidation on a longer
374 timescale or by the aqueous pathway for the formation of highly oxidized SOA in the ambient
375 air (Zhang et al., 2015).

376 Furthermore, the mass spectra of OA inside the chamber shows strong correlation with
377 the less-oxidized oxygenated organic aerosols (LO-OOA) derived from field measurements
378 in Beijing (Hu et al., 2016), which likely arose from oxidation of aromatic VOCs emitted
379 from vehicles (Peng et al., 2017). The correlation coefficient (R^2) initially was 0.88 and raised
380 to 0.99 sharply (Fig. 8B), indicating that the chamber well simulated the formation of
381 LO-OOA.

382 In our study, the secondary inorganic aerosols, i.e., sulfate, nitrate and ammonium, only
383 accounted for less than 10% of the coating materials on BC particles. This is consistent with
384 the previous studies showing that the concentration of organics is much larger than those of
385 sulfate and nitrate during the early stage of haze development in Beijing (Guo et al., 2014).
386 The low observed sulfate concentration in this study suggests that the gas phase formation of
387 sulfuric acid was unimportant under our experimental conditions. On the other hand, it has
388 been shown that the aqueous-phase reactions represent the dominate pathway for sulfate
389 formation in Beijing (Guo et al., 2010; Wang et al., 2016).

390 NO_2 has a higher reaction coefficient with the OH radical ($8 \times 10^{-12} cm^3 molecule^{-1} s^{-1}$)
391 than SO_2 (Zhang et al., 2015). Nitrate acid formed in the gas phase is transformed into nitrate
392 salts by the reaction with ammonia in the equilibrium process:





394 The equilibrium of this reaction is highly depended on ambient temperature and RH (Zheng
395 et al., 2008). In this study, chamber experiments were conducted in the afternoon with high
396 temperature and low RH (Table 1), which shifted the thermodynamic equilibrium to the gas
397 phase.

398 4.3 Hygroscopicity evolution

399 HTDMA measurement

400 The hygroscopic growth factors (HGF) of particles in each experiment were
401 continuously measured by the HTDMA system and corrected for the reference “dry”
402 diameters,

$$403 \quad HGF = \frac{D_{wet,t}/D_{m,t}}{D_{dry,0}/D_{m,0}} \quad (9)$$

404 where D_m is the mobility diameter of fresh or coated particles at dry condition, D_{dry} is the
405 mobility diameter of particles after experiencing a low humidity (below 30%) cycle in
406 HTDMA, and D_{wet} is the mobility diameter of particles after experiencing a high humidity
407 cycle (87%) in HTDMA.

408 Figure 9 shows the hygroscopicity variation of BC particles with the initial mobility
409 diameter ($D_{m,0}$) of 100 nm and 150 nm. The measured HGF of 0.999 - 1.004 for fresh BC
410 particles suggests high hydrophobicity, consistent with the previous studies (Khalizov et al.,
411 2009; Weingartner et al., 1997). After exposed to sunlight and ambient gaseous pollutants for
412 several hours, the HGF of these BC particles increased to 1.02-1.08 at the end of each
413 experiment. The HGF value varied with the total growth (ΔD_{me}) of BC particles, but was
414 constant at the same ΔD_{me} for different experiments (Fig.9). The final HGF values shown in
415 Figure 9 (1.02-1.08) were much lower than those in previous laboratory studies (Khalizov et
416 al., 2009; Tritscher et al., 2011) but similar to the low hygroscopic fraction in field
417 observations (Swietlicki et al., 2008), even for growth of particle size up to 90 nm in our
418 experiments.

419 The HGF is affected by many factors, e.g., the particle chemical composition and
420 morphology as well as RH (Qiu et al., 2012). The hygroscopicity of BC particles coated with
421 inorganic components, i.e., sulfuric acid (Khalizov et al., 2009), is significantly higher than
422 that coated by organic compounds (Tritscher et al., 2011). In this study, the major component
423 of the coating substance was LO-OOA with a O/C ratio about 0.5. The low oxygen content of
424 SOA coated on BC particles explains the low hygroscopicity (Jimenez et al., 2009),
425 indicating that coating of BC particles during the early stage haze development in Beijing
426 does not considerably increase the particle hygroscopicity.

427 The morphology of BC particles directly affects the HGF. As illustrated in Figure 10,
428 when the ΔD_{me} was 18 nm and 22 nm for 100 nm and 150 nm BC particles, respectively, the
429 HGF decreased slightly to about 0.99, suggesting that a thin layer of coatings on BC particles
430 decreased the particle diameter, even though a certain amount of water absorbed by BC
431 particles increased the particle mass. The surface tension of the water layer produced an



432 inward force on the “chain-like” branches of BC particles, leading to particle reconstruction,
433 and a more compact morphology. Such change was also identified in laboratory studies
434 (Weingartner et al., 1997; Tritscher et al., 2011; Qiu et al., 2012). In this study, the BC
435 particles became spherical when ΔD_{me} was 30 nm and 40 nm for particles with initial D_m of
436 100 nm and 150 nm, respectively (Peng et al., 2016). Therefore, when ΔD_{me} was large, the
437 HGF value was not influenced by reconstruction.

438 CCN measurements and κ closure

439 The CCN activation fraction (f_{CCN}) of BC particles at different supersaturation during two
440 typical experiments is illustrated in Figure 10. Fresh BC particles were not activated even at
441 very high supersaturation conditions (0.7%). With aging, f_{CCN} rapidly raised to nearly 100%
442 at high supersaturation (0.7% for experiment #4 and 0.6% for experiment #6). After several
443 hours, BC particles became CCN at lower supersaturation. The f_{CCN} at 0.4 supersaturation
444 (Fig. 11a in experiment #4) and 0.3 supersaturation (Fig. 11b in experiment #6) exceeded 50%
445 before the end of these two experiments, suggesting that aging increases the ability of BC
446 particles to become CCN (Wittbom et al., 2014) and a large amount of coatings results in
447 activation at lower supersaturation.

448 To further investigate the hygroscopicity of BC particles and combine the measurements
449 using HTDMA and CCN, we evaluated the hygroscopicity parameter, kappa (κ) (Petters and
450 Kreidenweis, 2007). The approximate relationship between the dry particle mass equivalent
451 diameter (D_{me}), the critical saturation ratio (S_c) and the apparent κ value of particles is
452 describe as:

$$453 \quad \kappa = \frac{4A^3}{27D_{me}\ln^2S_c} \quad (10)$$

454 where A is a parameter that includes several features of the solvent,

$$455 \quad A = \frac{4\sigma_{s/a}M_w}{RT\rho_w} \quad (11)$$

456 M_w is the molecular weight of water, ρ_w is the density of water, $\sigma_{s/a}$ is the surface tension of
457 the solution/air interface, R is the universal gas constant, and T is temperature.

458 In addition to the supersaturated condition, the κ theory also adopts the form for the
459 subsaturated condition, using the HGF from HTDMA measurement and RH:

$$460 \quad \frac{RH}{\exp\left(\frac{A}{D_{wet}}\right)} = \frac{HGF^3 - 1}{HGF^3 - (1 - \kappa)} \quad (12)$$

461 where D_{wet} is the wet diameter of particles.

462 The apparent κ values of BC particles calculated by HTDMA (κ_{HTDMA}) and CCN
463 (κ_{CCN}) are shown in Figure 10. The κ of fresh BC particles was near zero. With aging,
464 SOA coated on BC particles increased the κ_{HTDMA} and κ_{CCN} to approximately 0.04,
465 although the κ_{HTDMA} and κ_{CCN} exhibited difference features. A slightly higher κ_{CCN}
466 than κ_{HTDMA} at the beginning of aging was identified, attributed to reconstruction of BC



467 particles after humidified and underestimation of HGF and thus the κ_{HTDMA} value. Such a
468 difference between κ_{CCN} and κ_{HTDMA} was also observed in previous studies (Tritscher
469 et al., 2011; Martin et al., 2013). Nevertheless, the apparent κ values from two both methods
470 were comparable at the end of both experiments.

471 Assuming that a simple mixing rule is applicable to coated BC particles, the κ for
472 coating materials can be calculated based on the volume fraction of BC and SOA:

$$473 \quad \kappa = \sum_i \varepsilon_i \kappa_i \quad (13)$$

474 where ε_i represents the volume fraction of species i .

475 The κ values of the coating materials were 0.04 at the end of our experiments for both
476 CCN and HTDMA method, much lower than that of ambient aerosols in Beijing (Gunthe et
477 al., 2011; Wu et al., 2016) and those of SOA in previous chamber studies (Jimenez et al.,
478 2009; Tritscher et al., 2011; Martin et al., 2013). As discussed above, the coating substances
479 on BC particles were mainly SOA formed from photochemical oxidation. The κ of SOA
480 depends on the oxidation degree, which is correlated to the O/C ratio (Jimenez et al., 2009).
481 The O/C ratio of the coating SOA was 0.5 in our experiment, likely explaining the low κ with
482 LO-OOA.

483 5 Conclusions

484 In this paper, we present measurements of aging and hygroscopicity of BC particles in
485 Beijing using the QUALITY chamber. The unique two sub-chamber design facilitates the
486 evaluation of aging of BC particles under ambient conditions, by mimicking the ambient
487 gaseous concentrations without the presence of ambient aerosols. High UV transmission
488 efficiency (50-60%) and negligible wall loss of primary gaseous pollutants are shown for the
489 chamber performance. The validation experiments demonstrate little differences in the
490 primary gas concentrations, temperature, and RH between the chamber and the atmosphere,
491 suggesting that the chamber captures the evolution of ambient conditions. In addition, our
492 results show sulfuric acid production correlated with SO₂, indicating that the chamber well
493 simulates photochemical-driven formation of low volatile gaseous species by the hydroxyl
494 radical.

495 BC aging experiments were performed using the QUALITY chamber in Beijing. Fast
496 growth of BC particles (on average $26 \pm 11 \text{ nm h}^{-1}$) was observed, and SOA was identified as
497 the dominate component of the coating materials on BC particles, while inorganic species,
498 such as sulfate and nitrate, were unimportant under our experimental condition and timescale.

499 The HGF of BC particles exhibited a very low value (1.02-1.08) after several hours aging.
500 A slight decrease of HGF with a thin coating layer indicated reconstruction of BC particles
501 after humidified. Also, a very low kappa value (0.035) for BC particles at both subsaturation
502 and supersaturation conditions were found, with HTDMA and CCN measurements. Hence,
503 our results indicate that initial photochemical aging of BC particles does not appreciably alter
504 the particle hygroscopicity in Beijing.

505

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517 **References**

- 518 Bond, T. C., Doherty, S. J., Fahey, D. W., Forster, P. M., Berntsen, T., DeAngelo, B. J.,
519 Flanner, M. G., Ghan, S., Kärcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P. K.,
520 Sarofim, M. C., Schultz, M. G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S.,
521 Bellouin, N., Guttikunda, S. K., Hopke, P. K., Jacobson, M. Z., Kaiser, J. W., Klimont, Z.,
522 Lohmann, U., Schwarz, J. P., Shindell, D., Storelvmo, T., Warren, S. G., and Zender, C. S.:
523 Bounding the role of black carbon in the climate system: A scientific assessment, *Journal*
524 *of Geophysical Research: Atmospheres*, 118, 5380-5552, 10.1002/jgrd.50171, 2013.
- 525 Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P.,
526 Hildebrandt Ruiz, L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue,
527 N. M., Jayne, J. T., and Worsnop, D. R.: Elemental ratio measurements of organic
528 compounds using aerosol mass spectrometry: characterization, improved calibration, and
529 implications, *Atmos. Chem. Phys.*, 15, 253-272, doi:10.5194/acp-15-253-2015, 2015.
- 530 Claeys, M.: Formation of Secondary Organic Aerosols Through Photooxidation of Isoprene,
531 *Science*, 303, 1173-1176, doi:10.1126/science.1092805, 2004.
- 532 DeCarlo, P. F., Slowik, J. G., Worsnop, D. R., Davidovits, P., and Jimenez, J. L.: Particle
533 Morphology and Density Characterization by Combined Mobility and Aerodynamic
534 Diameter Measurements. Part 1: Theory, *Aerosol Sci. Tech.*, 38, 1185-1205,
535 10.1080/027868290903907, 2004.
- 536 DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C.,
537 Gonin, M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez, J. L.:
538 Field-deployable, high-resolution, time-of-flight aerosol mass spectrometer, *Anal. Chem.*,
539 78, 8281-8289, doi:10.1021/ac061249n, 2006.
- 540 DeCarlo, P. F., Ulbrich, I. M., Crouse, J., de Foy, B., Dunlea, E. J., Aiken, A. C., Knapp, D.,
541 Weinheimer, A. J., Campos, T., Wennberg, P. O., and Jimenez, J. L.: Investigation of the
542 sources and processing of organic aerosol over the Central Mexican Plateau from aircraft
543 measurements during MILAGRO, *Atmos. Chem. Phys.*, 10, 5257-5280,
544 doi:10.5194/acp-10-5257-2010, 2010.
- 545 Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud
546 droplets and aqueous particles (aqSOA): a review of laboratory, field and model studies,
547 *Atmos. Chem. Phys.*, 11, 11069-11102, doi:10.5194/acp-11-11069-2011, 2011.
- 548 Fortner, E. C., Zhao, J., and Zhang, R. Y.: Development of ion drift-chemical ionization mass
549 spectrometry, *Anal. Chem.*, 76, 5436-5440, doi:10.1021/ac0493222, 2004.
- 550 Gunthe, S. S., Rose, D., Su, H., Garland, R. M., Achtert, P., Nowak, A., Wiedensohler, A.,
551 Kuwata, M., Takegawa, N., Kondo, Y., Hu, M., Shao, M., Zhu, T., Andreae, M. O., and
552 Pöschl, U.: Cloud condensation nuclei (CCN) from fresh and aged air pollution in the
553 megacity region of Beijing, *Atmos. Chem. Phys.*, 11, 11023-11039,
554 doi:10.5194/acp-11-11023-2011, 2011.
- 555 Guo, S., Hu, M., Lin, Y., Gomez-Hernandez, M., Zamora, M. L., Peng, J. F., Collins, D. R.,
556 and Zhang, R. Y.: OH-Initiated Oxidation of m-Xylene on Black Carbon Aging, *Environ.*
557 *Sci. Technol.*, 50, 8605-8612, doi:10.1021/acs.est.6b01272, 2016.
- 558 Guo, S., Hu, M., Wang, Z. B., Slanina, J., and Zhao, Y. L.: Size-resolved aerosol
559 water-soluble ionic compositions in the summer of Beijing: implication of regional
560 secondary formation, *Atmos. Chem. Phys.*, 10, 947-959, 2010.



- 561 Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M.,
562 Zeng, L., Molina, M. J., and Zhang, R.: Elucidating severe urban haze formation in China,
563 *P. Natl. Acad. Sci. USA*, 111, 17373-17378, doi:10.1073/pnas.1419604111, 2014.
- 564 He, L.-Y., X.-F. Huang, L. Xue, M. Hu, Y. Lin, J. Zheng, R. Zhang, and Y.-H. Zhang:
565 Submicron aerosol analysis and organic source apportionment in an urban atmosphere in
566 Pearl River Delta of China using high-resolution aerosol mass spectrometry, *J. Geophys.*
567 *Res.* 116, D12304, doi:10.1029/2010JD014566, 2011.
- 568 Hu, M., Peng, J., Sun, K., Yue, D., Guo, S., Wiedensohler, A., and Wu, Z.: Estimation of
569 Size-Resolved Ambient Particle Density Based on the Measurement of Aerosol Number,
570 Mass, and Chemical Size Distributions in the Winter in Beijing, *Environ. Sci. Technol.*,
571 46, 9941-9947, doi:10.1021/es204073t, 2012.
- 572 Hu, W., Hu, M., Hu, W., Jimenez, J. L., Yuan, B., Chen, W., Wang, M., Wu, Y., Chen, C.,
573 Wang, Z., Peng, J., Zeng, L., and Shao, M.: Chemical composition, sources and aging
574 process of sub-micron aerosols in Beijing: contrast between summer and winter, *J.*
575 *Geophys. Res.- Atmos.*, 121, 1955-1977 doi:10.1002/2015jd024020, 2016.
- 576 Huang, X. F., He, L. Y., Hu, M., Canagaratna, M. R., Sun, Y., Zhang, Q., Zhu, T., Xue, L.,
577 Zeng, L. W., Liu, X. G., Zhang, Y. H., Jayne, J. T., Ng, N. L., and Worsnop, D. R.: Highly
578 time-resolved chemical characterization of atmospheric submicron particles during 2008
579 Beijing Olympic Games using an Aerodyne High-Resolution Aerosol Mass Spectrometer,
580 *Atmos. Chem. Phys.*, 10, 8933-8945, doi:10.5194/acp-10-8933-2010, 2010.
- 581 Intergovernmental Panel on Climate Change (IPCC), *Climate Change 2013: The Physical*
582 *Science Basis*, Cambridge Univ. Press, Cambridge, UK, 2013.
- 583 Jacobson, M. Z.: Strong radiative heating due to the mixing state of black carbon in
584 atmospheric aerosols, *Nature*, 409, 695-697, doi:10.1038/35055518, 2001.
- 585 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H.,
586 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I.
587 M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A.,
588 Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J.,
589 Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J.,
590 Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K.,
591 Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo,
592 D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J.
593 Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C.,
594 Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E.,
595 Baltensperger, U., and Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere,
596 *Science*, 326, 1525-1529, doi:10.1126/science.1180353, 2009.
- 597 Kalberer, M.: Identification of Polymers as Major Components of Atmospheric Organic
598 Aerosols, *Science*, 303, 1659-1662, doi:10.1126/science.1092185, 2004.
- 599 Khalizov, A. F., Xue, H., and Zhang, R.: Enhanced light absorption and scattering by carbon
600 soot aerosols internally mixed with sulfuric acid, *J. Phys. Chem.*, 113, 1066-1074,
601 doi:10.1021/jp807531n, 2009.
- 602 Khalizov, A. F., Zhang, R., Zhang, D., Xue, H., Pagels, J., and McMurry, P. H.: Formation of
603 highly hygroscopic soot aerosols upon internal mixing with sulfuric acid vapor, *J.*
604 *Geophys. Res.*, 114, doi:10.1029/2008jd010595, 2009.



- 605 Khalizov, A. F., Lin, Y., Qiu, C., Guo, S., Collins, D., and Zhang, R.: Role of OH-Initiated
606 Oxidation of Isoprene in Aging of Combustion Soot, *Environ. Sci. Technol.*, 47,
607 2254-2263, doi:10.1021/es3045339, 2013.
- 608 Kuwata, M., Kondo, Y., Mochida, M., Takegawa, N., and Kawamura, K.: Dependence of
609 CCN activity of less volatile particles on the amount of coating observed in Tokyo, *J.*
610 *Geophys. Res.*, 112, doi:10.1029/2006jd007758, 2007.
- 611 Liu, S., Aiken, A. C., Gorkowski, K., Dubey, M. K., Cappa, C. D., Williams, L. R., Herndon,
612 S. C., Massoli, P., Fortner, E. C., Chhabra, P. S., Brooks, W. A., Onasch, T. B., Jayne, J. T.,
613 Worsnop, D. R., China, S., Sharma, N., Mazzoleni, C., Xu, L., Ng, N. L., Liu, D., Allan, J.
614 D., Lee, J. D., Fleming, Z. L., Mohr, C., Zotter, P., Szidat, S., and Prevot, A. S. H.:
615 Enhanced light absorption by mixed source black and brown carbon particles in UK
616 winter, *Nat. Commun.*, 6, Artn 8435, doi:10.1038/Ncomms9435, 2015.
- 617 Ma, Y., Brooks, S. D., Vidaurre, G., Khalizov, A. F., Wang, L., and Zhang, R.: Rapid
618 modification of cloud-nucleating ability of aerosols by biogenic emissions, *Geophys. Res.*
619 *Lett.*, 40, 6293-6297, doi:10.1002/2013gl057895, 2013.
- 620 Martin, M., Tritscher, T., Jurányi, Z., Heringa, M. F., Sierau, B., Weingartner, E., Chirico, R.,
621 Gysel, M., Prévôt, A. S. H., Baltensperger, U., and Lohmann, U.: Hygroscopic properties
622 of fresh and aged wood burning particles, *J. Aerosol Sci.*, 56, 15-29,
623 doi:10.1016/j.jaerosci.2012.08.006, 2013.
- 624 Moffet, R. C., and Prather, K. A.: In-situ measurements of the mixing state and optical
625 properties of soot with implications for radiative forcing estimates, *P. Natl. Acad. Sci.*
626 *USA*, 106, 11872-11877, doi:10.1073/pnas.0900040106, 2009.
- 627 Pagels, J., McMurry, P.H., Khalizov, A.F., and Zhang, R.: Processing of soot by controlled
628 sulphuric acid and water condensation—Mass and mobility relationship, *Aerosol Sci.*
629 *Tech.* 43, 629–640, 2009.
- 630 Peng, J., Hu, M., Du, Z., Wang, Y., Zheng, J., Zhang, W., Yang, Y., Qin, Y., Zheng, R., Xiao,
631 Y., Wu, Y., Lu, S., Wu, Z., Guo, S., Mao, H., and Shuai, S.: Gasoline aromatic: a critical
632 determinant of urban secondary organic aerosol formation, *Atmos. Chem. Phys. Discuss.*,
633 doi:10.5194/acp-2017-254, in review, 2017.
- 634 Peng, J. F., Hu, M., Guo, S., Du, Z. F., Zheng, J., Shang, D. J., Zamora, M. L., Zeng, L. M.,
635 Shao, M., Wu, Y. S., Zheng, J., Wang, Y., Glen, C. R., Collins, D. R., Molina, M. J., and
636 Zhang, R. Y.: Markedly enhanced absorption and direct radiative forcing of black carbon
637 under polluted urban environments, *P. Natl. Acad. Sci. USA*, 113, 4266-4271,
638 doi:10.1073/pnas.1602310113, 2016.
- 639 Peng, J. F., Hu, M., Wang, Z. B., Huang, X. F., Kumar, P., Wu, Z. J., Guo, S., Yue, D. L.,
640 Shang, D. J., Zheng, Z., and He, L. Y.: Submicron aerosols at thirteen diversified sites in
641 China: size distribution, new particle formation and corresponding contribution to cloud
642 condensation nuclei production, *Atmos. Chem. Phys.*, 14, 10249-10265,
643 doi:10.5194/acp-14-10249-2014, 2014.
- 644 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic
645 growth and cloud condensation nucleus activity, *Atmos. Chem. Phys.*, 7, 1961-1971,
646 2007.
- 647 Qiu, C., Khalizov, A. F., and Zhang, R. Y.: Soot Aging from OH-Initiated Oxidation of
648 Toluene, *Environ. Sci. Technol.*, 46, 9464-9472, doi:10.1021/Es301883y, 2012.



- 649 Qiu, C., and Zhang, R.: Multiphase chemistry of atmospheric amines, *Phys. Chem. Chem.*
650 *Phys.* 15, doi:10.1039/C3CP43446J, 5738-5752, 2013.
- 651 Reed, R. C.: Observations of secondary organic aerosol production and soot aging under
652 atmospheric conditions using a novel new environmental aerosol chamber, PhD
653 dissertation, Texas A&M Univ, College Station, TX., 2010.
- 654 Ruehl, C. R., Davies, J. F., and Wilson, K. R.: An interfacial mechanism for cloud droplet
655 formation on organic aerosols, *Science*, 351, 1447-1450, doi:10.1126/science.aad4889,
656 2016.
- 657 Saathoff, H., Naumann, K. H., Schnaiter, M., Schöck, W., Möhler, O., Schurath, U.,
658 Weingartner, E., Gysel, M., and Baltensperger, U.: Coating of soot and (NH₄)₂SO₄
659 particles by ozonolysis products of α -pinene, *J. Aerosol Sci.*, 34, 1297-1321,
660 doi:10.1016/s0021-8502(03)00364-1, 2003.
- 661 Santoro, R. J., Semerjian, H. G., and Dobbins, R. A.: Soot Particle Measurements in
662 Diffusion Flames, *Combust. Flame*, 51, 203-218, 1983.
- 663 Swietlicki, E., Hansson, H. C., HÄMeri, K., Svenningsson, B., Massling, A., McFiggans, G.,
664 McMurry, P. H., PetÄJÄ, T., Tunved, P., Gysel, M., Topping, D., Weingartner, E.,
665 Baltensperger, U., Rissler, J., Wiedensohler, A., and Kulmala, M.: Hygroscopic properties
666 of submicrometer atmospheric aerosol particles measured with H-TDMA instruments in
667 various environments—a review, *Tellus B*, 60, 432-469,
668 doi:10.1111/j.1600-0889.2008.00350.x, 2008.
- 669 Tritscher, T., Jurányi, Z., Martin, M., Chirico, R., Gysel, M., Heringa, M. F., DeCarlo, P. F.,
670 Sierau, B., Prévôt, A. S. H., Weingartner, E., and Baltensperger, U.: Changes of
671 hygroscopicity and morphology during ageing of diesel soot, *Environ. Res. Lett.*, 6,
672 034026, doi:10.1088/1748-9326/6/3/034026, 2011.
- 673 Wang, G. H., Zhang, R. Y., Gomez, M. E., Yang, L. X., Zamora, M. L., Hu, M., Lin, Y., Peng,
674 J. F., Guo, S., Meng, J. J., Li, J. J., Cheng, C. L., Hu, T. F., Ren, Y. Q., Wang, Y. S., Gao, J.,
675 Cao, J. J., An, Z. S., Zhou, W. J., Li, G. H., Wang, J. Y., Tian, P. F., Marrero-Ortiz, W.,
676 Secretst, J., Du, Z. F., Zheng, J., Shang, D. J., Zeng, L. M., Shao, M., Wang, W. G., Huang,
677 Y., Wang, Y., Zhu, Y. J., Li, Y. X., Hu, J. X., Pan, B., Cai, L., Cheng, Y. T., Ji, Y. M.,
678 Zhang, F., Rosenfeld, D., Liss, P. S., Duce, R. A., Kolb, C. E., and Molina, M. J.:
679 Persistent sulfate formation from London Fog to Chinese haze, *P. Natl. Acad. Sci. USA*,
680 113, 13630-13635, doi:10.1073/pnas.1616540113, 2016.
- 681 Wang, L., Khalizov, A. F., Zheng, J., Xu, W., Ma, Y., Lal, V., and Zhang, R. Y.: Atmospheric
682 nanoparticles formed from heterogeneous reactions of organics, *Nat. Geosci.*, 3, 238-242,
683 doi:10.1038/NGEO778, 2010.
- 684 Wang, Y., Wan, Q., Meng, W., Liao, F., Tan, H., and Zhang, R.: Long-term impacts of
685 aerosols on precipitation and lightning over the Pearl River Delta megacity area in China,
686 *Atmos. Chem. Phys.* 11, 12421-12436, 2011.
- 687 Wang, Y., Khalizov, A., Levy, M., and Zhang, R.: Light absorbing aerosols and their
688 atmospheric impacts, *Atmos. Environ.* 81, 713-715, doi:10.1016/j.atmosenv.2013.09.034,
689 2013.
- 690 Weingartner, E., Burtscher, H., and Baltensperger, U.: Hygroscopic properties of carbon and
691 diesel soot particles, *Atmos. Environ.*, 31, 2311-2327,
692 doi:10.1016/S1352-2310(97)00023-X, 1997.



- 693 Wittbom, C., Eriksson, A. C., Rissler, J., Carlsson, J. E., Roldin, P., Nordin, E. Z., Nilsson, P.
694 T., Swietlicki, E., Pagels, J. H., and Svenningsson, B.: Cloud droplet activity changes of
695 soot aerosol upon smog chamber ageing, *Atmos. Chem. Phys.*, 14, 9831-9854,
696 doi:10.5194/acp-14-9831-2014, 2014.
- 697 Wu, Z. J., Zheng, J., Shang, D. J., Du, Z. F., Wu, Y. S., Zeng, L. M., Wiedensohler, A., and
698 Hu, M.: Particle hygroscopicity and its link to chemical composition in the urban
699 atmosphere of Beijing, China, during summertime, *Atmos. Chem. Phys.*, 16, 1123-1138,
700 doi:10.5194/acp-16-1123-2016, 2016.
- 701 Xue, H., Khalizov, A. F., Wang, L., Zheng, J., and Zhang, R.: Effects of dicarboxylic acid
702 coating on the optical properties of soot, *Phys. Chem. Chem. Phys.*, 11, 7865-7875,
703 doi:10.1039/b700001a, 2009a.
- 704 Xue, H., Khalizov, A. F., and Zhang, R.: Effects of coating of dicarboxylic acids on the
705 mass-mobility relationship of soot particles, *Environ. Sci. Technol.* 43, 2787-2792, doi:
706 10.1021/es803287v, 2009b.
- 707 Yuan, T., Li, Z., Zhang, R., and Fan, J.: Increase of cloud droplet size with aerosol optical
708 depth: An observation and modeling study, *J. Geophys. Res.*, 113, D04201,
709 doi:10.1029/2007JD008632, 2008.
- 710 Zhang, D., and Zhang R.: Laboratory investigation of heterogeneous interaction of sulfuric
711 acid with soot, *Environ. Sci. Technol.* 39, 5722-5727, doi:10.1021/es050372d, 2005.
- 712 Zhang, R., Khalizov, A. F., Pagels, J., Zhang, D., Xue, H., and McMurry, P. H.: Variability in
713 morphology, hygroscopicity, and optical properties of soot aerosols during atmospheric
714 processing, *P. Natl. Acad. Sci. USA*, 105, 10291-10296, doi:10.1073/pnas.0804860105,
715 2008.
- 716 Zhang, R. Y., Suh, I., Zhao, J., Zhang, D., Fortner, E. C., Tie, X. X., Molina, L. T., and
717 Molina, M. J.: Atmospheric new particle formation enhanced by organic acids, *Science*,
718 304, 1487-1490, doi:10.1126/science.1095139, 2004.
- 719 Zhang, R., Khalizov, A.F., Wang, L., Hu, M., Xu, W.: Nucleation and growth of nanoparticles
720 in the atmosphere, *Chem. Rev.*, 112, 1957-2011, doi:10.1021/cr2001756, 2012.
- 721 Zhang, R. Y., Wang, G. H., Guo, S., Zarnora, M. L., Ying, Q., Lin, Y., Wang, W. G., Hu, M.,
722 and Wang, Y.: Formation of Urban Fine Particulate Matter, *Chem. Rev.*, 115, 3803-3855,
723 doi:10.1021/acs.chemrev.5b00067, 2015.
- 724 Zhao, J., Levitt, N. P., Zhang, R., and Chen, J.: Heterogeneous reactions of methylglyoxal in
725 acidic media: Implications for secondary organic aerosol formation, *Environ. Sci.*
726 *Technol.*, 40, 7682-7687, doi:10.1021/es060610k, 2006.
- 727 Zhao, J., Khalizov, A.F., Zhang, R., and McGraw, R.: Hydrogen bonding interaction of
728 molecular complexes and clusters of aerosol nucleation precursors, *J. Phys. Chem.*, 113,
729 680-689, doi:10.1021/jp806693r, 2009.
- 730 Zhao, P. S., Dong, F., He, D., Zhao, X. J., Zhang, X. L., Zhang, W. Z., Yao, Q., and Liu, H. Y.:
731 Characteristics of concentrations and chemical compositions for PM_{2.5} in the region of
732 Beijing, Tianjin, and Hebei, China, *Atmos. Chem. Phys.*, 13, 4631-4644,
733 doi:10.5194/acp-13-4631-2013, 2013.
- 734 Zheng, J., Zhang, R., Fortner, E. C., Volkamer, R. M., Molina, L., Aiken, A. C., Jimenez, J. L.,
735 Gaeggeler, K., Dommen, J., Dusanter, S., Stevens, P.S., and Tie, X.: Measurements of
736 HNO₃ and N₂O₅ using ion drift - chemical ionization mass spectrometry during the



- 737 MCMA - 2006 campaign, Atmos. Chem. Phys., 8, 6823–6838, 2008.
- 738 Zheng, J., Khalizov, A., Wang, L., and Zhang, R. Y.: Atmospheric Pressure-Ion Drift
- 739 Chemical Ionization Mass Spectrometry for Detection of Trace Gas Species, Anal. Chem.,
- 740 82, 7302-7308, doi:10.1021/ac101253n, 2010.



Table 1. Summary of ambient conditions for BC aging experiments conducted in Beijing. The $PM_{2.5}$, gas concentrations and meteorological condition were averaged from the entire experimental time. D_m and T represent the mobility diameter and temperature, respectively.

No.	Date	Time		Initial D_m (nm)	$PM_{2.5}$ ($\mu g\ m^{-3}$)	$J(O^1D)$ (10^{-6})	Gas Concentration (ppb)				Meteorological Conditions		
		Start	End				Toluene	Xylene	SO ₂	NO _x	O ₃	T (°C)	RH (%)
#1	Aug. 18 th	12:41	14:47	95	43	19	0.49	0.13	1.7	8.8	56	35	27
#2	Aug. 22 nd	13:32	16:51	96	69	3.7	3.41	0.78	2	36.1	26	26	69
#3	Sep. 7 th	12:40	14:52	97	12	17.5	0.71	0.17	2.7	10.2	75	30	35
#4	Sep. 9 th	13:13	17:06	97	40	6.3	0.77	0.21	4	11	68	26	50
#5	Sep. 1 st	13:19	16:13	147	27	11.6	0.76	0.19	4.6	19.9	83	31	33
#6	Sep. 11 st	13:50	17:25	147	57	6.1	1.57	0.39	6.7	17.1	92	29	42
#7	Sep. 21 st	15:31	17:41	146	30	2.1	0.75	0.29	2	10.6	90	28	37
#8	Aug. 24 th	11:37	16:06	216	8.8	21.1	0.98	0.3	1.7	15.6	57	36	25
#9	Sep. 5 th	14:06	16:44	220	12	8.3	0.45	0.2	2.2	14.6	54	29	35
#10	Oct. 17 th	12:54	17:13	224	57	3.2	-	-	13.8	41	34	18	30



Table 2. Summary of particle properties for BC aging experiments conducted in Beijing. D_m , ρ_{eff} and D_{me} represent the mobility diameter, effective density and mass equivalent diameter, respectively.

No.	Date	D_m		ρ_{eff}		D_{me}			
		Initial (nm)	Final (nm)	Initial (g cm^{-3})	Final (g cm^{-3})	Initial (nm)	Final (nm)	ΔD_{me} (nm)	Growth Rate ^a (nm h^{-1})
#1	Aug. 18 th	95	157	0.50	1.35	62	162	100	47
#2	Aug. 22 nd	96	129	0.46	1.31	61	126	65	20
#3	Sep. 7 th	97	147	0.45	1.25	62	142	80	36
#4	Sep. 9 th	97	136	0.43	1.30	61	133	62	19
#5	Sep. 1 st	147	170	0.34	1.36	85	168	83	29
#6	Sep. 11 st	147	162	0.34	1.34	84	159	75	20
#7	Sep. 21 st	146	132	0.34	1.05	84	116	32	15
#8	Aug. 24 th	216	272	0.32	1.37	123	275	152	34
#9	Sep. 5 th	220	202	0.25	1.33	114	197	83	31
#10	Oct. 17 th	224	224	0.24	0.52	117	157	40	11
		Average						77 ± 33	26 ± 11

^a The growth rate is calculated using the data between 12:00 and 17:00 for each experiment.

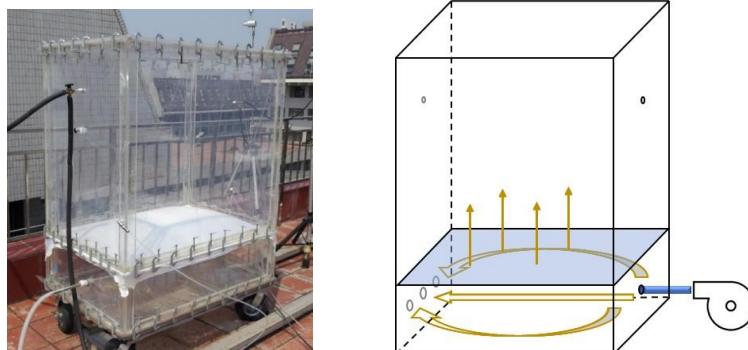


Figure 1. Photo (left) and schematic (right) of the quasi-atmospheric aerosol evolution study (QUALITY) chamber

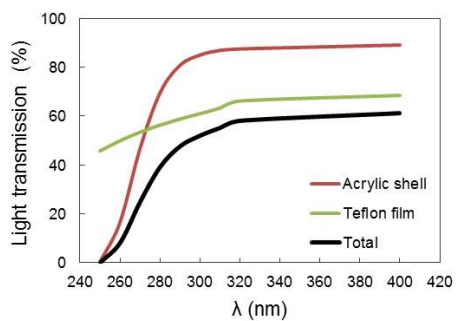


Figure 2. Light transmission spectra of the PFA Teflon film (yellow line), the acrylite shell (red line), and their total transmission in the UV range (black line).

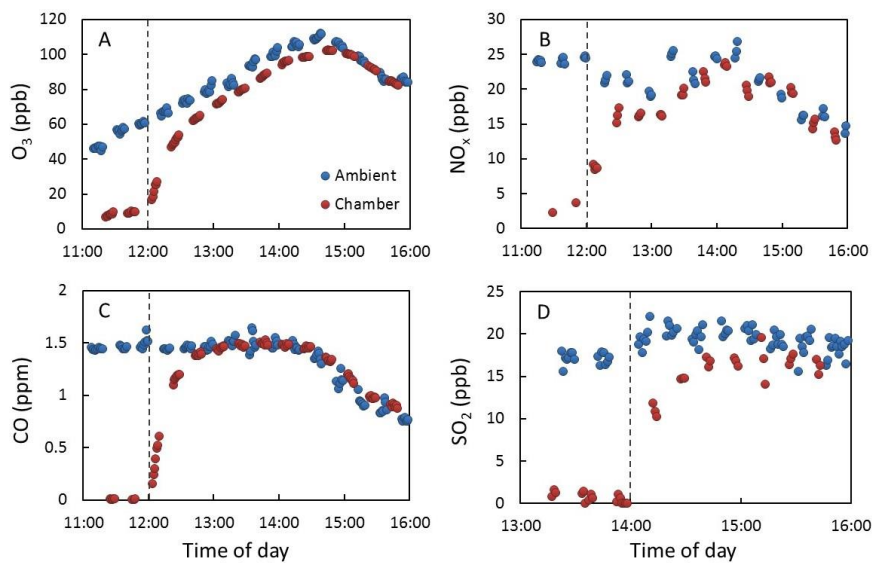


Figure 3. Concentrations of O₃ (A), NO_x (B), CO (C), and SO₂ (D) measured inside the QUALITY chamber (Red circles) and in the ambient air (blue circles). The vertical dashed lines denote the time when the ambient air started to be pulled through the bottom flow chamber and the ambient gases began to exchange into the upper reaction chamber.

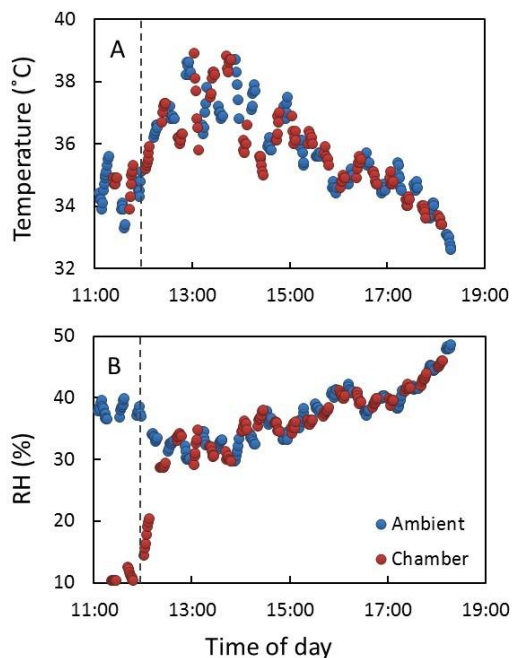


Figure 4. Temperature (A) and RH (B) measured inside the QUALITY chamber (Red circles) and in the ambient air (blue circles). The temperature and RH were measured by probes placed in the sampling tube adjacent to the chamber. The vertical dashed lines denote the time when the ambient air started to be pulled through the bottom flow chamber and the ambient gases began to exchange into the upper reaction chamber.

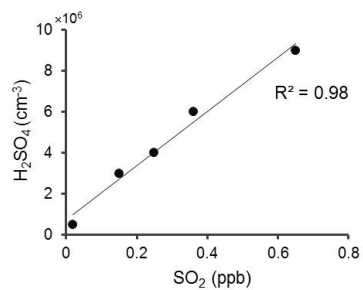


Figure 5. Sulfuric acid concentration as function of SO₂ concentration inside the chamber

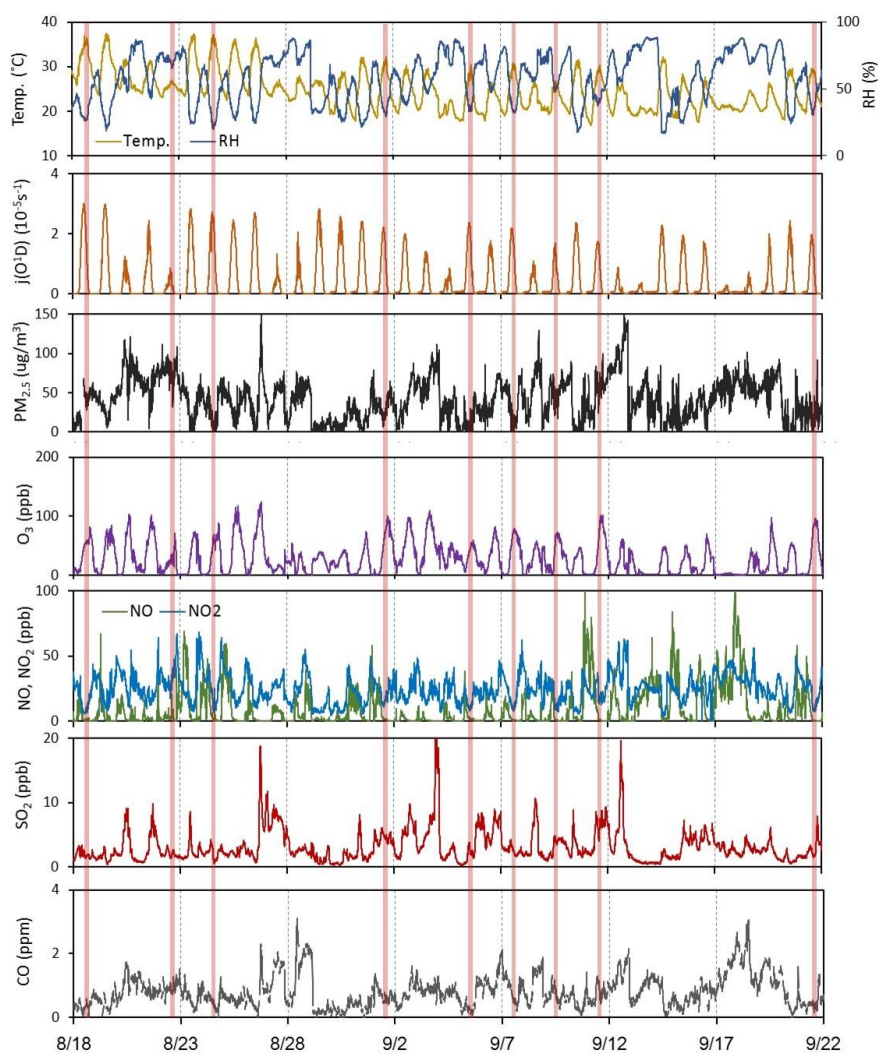


Figure 6. Time series of ambient pollutant concentrations and meteorological parameters during the experimental period in Beijing. Red bars indicate periods when BC aging experiments were conducted.

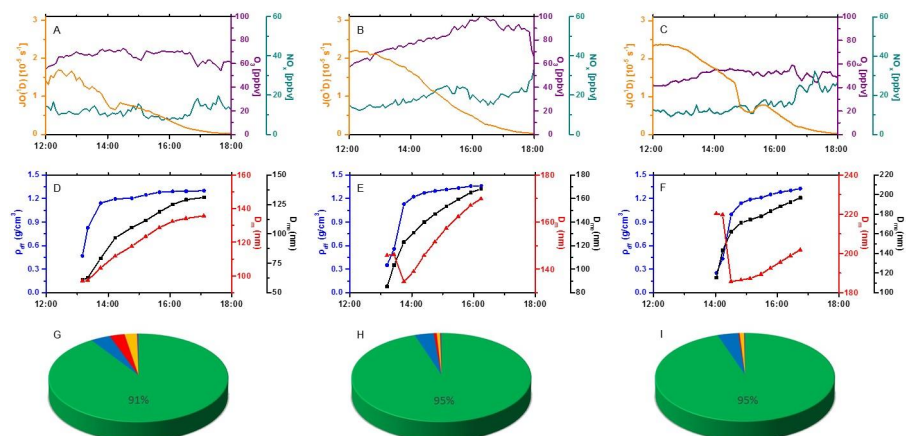


Figure 7. Ambient condition (A, B, C), changes of diameter and density of BC particle (D, E, F), and the chemical composition of coating materials (G, H, I) during three typical aging experiments. A and D correspond to experiment #4; B and E correspond to experiment #5, and C and E correspond to experiment #9. D_m is the peak mobility diameter of BC particles, ρ_{eff} is the best-fit effective density of BC particles with the mobility diameter D_m , D_{me} is the mass equivalent diameter of BC particles, and $J(O^1D)$ represents the measured photolysis rate constant for $O(^1D)$. The colors of green, blue, red, yellow and purple in the pie charts represent organics, nitrate, sulfate, ammonium, and chlorine, respectively. The numbers in figure G, H and I are the mass fraction of organics.

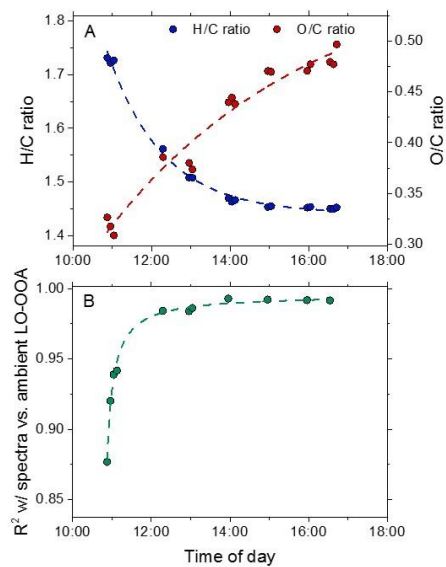


Figure 8. The evolution of organic aerosols inside the chamber during an aging experiment (#8) in Beijing. (A) the H/C and O/C ratios of organics on aged BC particles; (b) the correlation coefficients (R^2) between the evolving total OA spectra in chamber experiment and the LO-OOA spectra derived from the Beijing field data set.

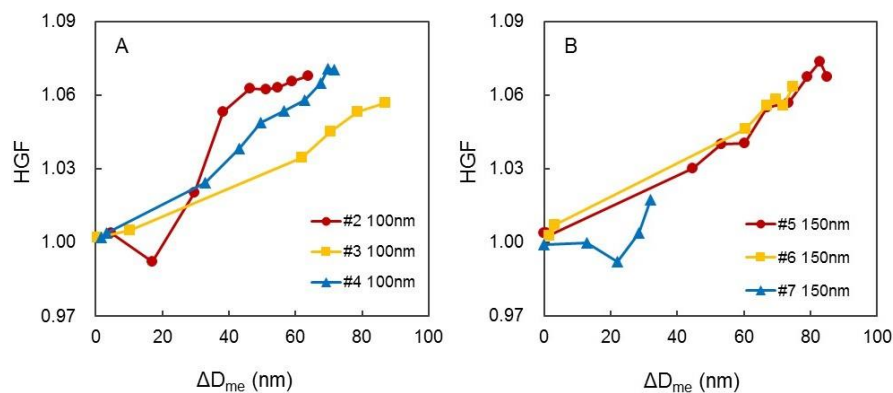


Figure 9. Evolution of hygroscopic growth factors (HGF) of BC particles during aging as a function of ΔD_{me} . (A) three experiments with 100 nm BC particle; (B) three experiments with 150 nm BC particle. Different colors in each figure represent different experiments. Hygroscopicity measurement is not available for experiment #1.

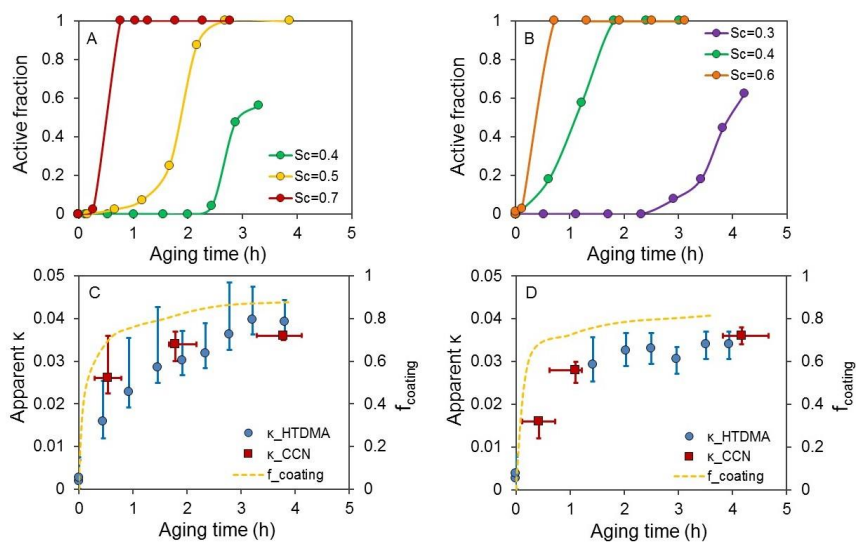


Figure 10. The active fraction of BC particles under diversified supersaturation (A, B) and the closure of apparent κ for BC particles with initial diameter (C, D) during aging in two typical experiments. A, C are the results from experiment #4 with 100nm BC particles and B, D represents experiment #6 with 150nm BC particles. Red square and blue circle in C and D represent the apparent κ calculated using CCN counter data and HTDMA data, respectively. Yellow slash line represents the fraction of coating materials on BC particles. The error bars of κ_{CCN} and κ_{HTDMA} represent the uncertainty in the calculation.