

1 **Aging and hygroscopicity variation of black carbon particles in Beijing measured by a**
2 **quasi-atmospheric aerosol evolution study (QUALITY) chamber**

3 Jianfei Peng^{1, 2*}, Min Hu^{1, 3*}, Song Guo^{1, 2}, Zhuofei Du¹, Dongjie Shang¹, Jing Zheng¹, Jun
4 Zheng², Limin Zeng¹, Min Shao¹, Yusheng Wu¹, Don Collins², Renyi Zhang^{1, 2*}

5 ¹State Key Joint Laboratory of Environmental Simulation and Pollution Control, College of
6 Environmental Sciences and Engineering, Peking University, Beijing, China 100871

7 ²Department of Atmospheric Sciences, Texas A&M University, College Station, Texas, USA
8 77843

9 ³Beijing Innovation Center for Engineering Sciences and Advanced Technology, Peking
10 University, Beijing, China 100871

11 *To whom correspondence should be addressed. E-mail: pengjianfeipku@gmail.com (Jianfei
12 Peng); minhu@pku.edu.cn (Min Hu); renyi-zhang@geos.tamu.edu (Renyi Zhang)

13 **Abstract.** Measurements of aging and hygroscopicity variation of black carbon (BC)
14 particles in Beijing were conducted using a 1.2 m³ quasi-atmospheric aerosol evolution study
15 (QUALITY) chamber, which consisted of a bottom flow chamber where ambient air was
16 pulled through continuously and an upper reaction chamber where aging of BC particles
17 occurred. Within the reaction chamber, transmission of the solar ultraviolet irradiation was
18 approximately 50% - 60%, wall loss of primary gaseous pollutants was negligible compared
19 with the replenish rate by gas exchange, and BC exhibited a half-lifetime about 3-7 hours.
20 Typically, equilibrium for the primary gases, temperature, and relative humidity between the
21 reaction chamber and ambient air was established within 1 hour. Rapid growth of BC
22 particles was observed, with an average total growth of 77±33nm and average growth rate of
23 26±11 nm h⁻¹. Secondary organic aerosols (SOA) accounted for more than 90% of the coating
24 mass. The O/C ratio of SOA was 0.5, lower than the ambient level. The hygroscopic growth
25 factor of BC particles decreased slightly with an initial thin coating layer because of BC
26 reconstruction, but subsequently increased to 1.06-1.08 upon further aging. The κ (kappa)
27 values for BC particles and coating materials were calculated as 0.035 and 0.040 at the
28 subsaturation and supersaturation conditions, respectively, indicating low hygroscopicity of
29 coated SOA on BC particles. Hence, our results indicate that initial photochemical aging of
30 BC particles does not appreciably alter the particle hygroscopicity in Beijing.

31

32 1 Introduction

33 Atmospheric aerosols undergo continuous and complicated transformation during their
34 residence time in the atmosphere. The aging of aerosols is likely resulted from both physical
35 (i.e., coagulation, condensation, equilibrium partitioning, and evaporation) and chemical (i.e.,
36 photochemical gas-phase oxidation and multi-phase reactions) processes (Zhao et al., 2006;
37 Qiu et al., 2013; Zhang et al., 2015). Also, there are typically large variations in the particle
38 properties (i.e., size, mass, chemical composition, morphology, and optical and hygroscopic
39 parameters) during aging, significantly influencing the aerosol impacts on visibility, human
40 health, weather, and climate (Jacobson, 2001; Guo et al., 2014). A better understanding of the
41 aging process of aerosols in the atmosphere is critical in atmospheric and climate research.

42 For example, the scientific interest in the climate effects of black carbon (BC) has
43 remained, since BC is the strongest absorber of visible solar radiation (Wang et al., 2013). BC
44 solar absorption represents a central issue in climate change research, since the synthesis of
45 satellite, in situ, and ground observations shows that the global solar absorption (direct
46 radiative forcing or DRF) by BC is as much as $0.9 \text{ W}\cdot\text{m}^{-2}$, second only to that of CO_2
47 (Jacobson, 2001; Bond et al., 2013; IPCC, 2013). BC also represents an important component
48 of air pollution for large parts of the world (Zhang et al., 2015). The properties of BC are
49 considerably modified during aging, including the size, mass, morphology, and optical and
50 hygroscopic parameters (Khalizov et al., 2009; Xue et al., 2009a). Enhanced light absorption
51 of BC particles during aging not only contributes to atmospheric stabilization and
52 exacerbation of haze formation, but also imposes large positive radiative forcing on climate
53 (Peng et al., 2016). Furthermore, the variation in hygroscopicity during aging also regulates
54 the lifetimes of BC particles. Hygroscopic particles serve efficiently as cloud condensed
55 nuclei (CCN), affecting the formation, longevity and albedo of clouds (Yuan et al., 2008;
56 Wang et al., 2011). Also, deposition of BC particles, via in-cloud scavenging and wet
57 deposition, depends highly on the particle hygroscopicity (Bond et al., 2013). In addition, the
58 hygroscopicity also affects the aqueous-phase reactions of atmospheric pollutants (Ervens et
59 al., 2011; Wang et al., 2016). Previous studies using hygroscopic tandem differential mobility
60 analyzer (H-TDMA) instruments have shown that coating of hydrophilic materials
61 significantly increases the hygroscopic growth factor of BC particles (Saathoff et al., 2003;
62 Khalizov et al., 2009; Guo et al., 2016). The ability of BC particles to form CCN also
63 enhanced after coating of hydrophilic materials (Kuwata et al., 2007; Tritscher et al., 2011;
64 Ma et al., 2013; Wittbom et al., 2014). The activation supersaturation depends on the particle
65 size, hygroscopicity of coating materials, and the coating thickness (Ma et al., 2013). The
66 coating materials in the previous experiments include sulfuric acid (Zhang and Zhang, 2005;
67 Khalizov et al., 2009), oxidation products from biogenic and anthropogenic hydrocarbon
68 species (Saathoff et al., 2003; Ma et al., 2013; Khalizov et al., 2013), or secondary organic
69 aerosols (SOA) from single emission source (Tritscher et al., 2011). However, there still exist
70 uncertainties for parametrization of the BC lifetime in atmospheric models, because of
71 insufficient constraints on the hydrophobic to hydrophilic conversion of BC particles under
72 variable ambient conditions.

73 Atmospheric field measurements have been performed to evaluate aging of particles on
74 different platforms, e.g., ground, aircraft, and cruise (Moffet and Prather, 2009; DeCarlo et al.,

75 2010; Peng et al., 2014; Liu et al., 2015) and over different spatial scales (intensive
76 campaigns or long-term measurements). Typically, a wide variety of state-of art instruments
77 are employed to characterize the changes of the chemical and physical properties of aerosols.
78 On the other hand, field measurements at fixed sites are affected by transport, local emissions,
79 and chemistry, and quantification of the particle parameters during aging involves complex
80 decoupling of the various processes (Peng et al., 2016). In particular, it is challenging to
81 isolate the chemical processes from those related to meteorology (i.e., transport and mixing)
82 and emissions.

83 The methods of environmental chambers or reactors have been widely employed in
84 atmospheric chemistry research, including photochemical oxidation of volatile organic
85 compounds (VOCs) (Zhang et al., 2000), formation and growth of aerosols (Claeys, 2004;
86 Kalberer, 2004), nucleation of nanoparticles (Zhang et al., 2004; Wang et al., 2010; Zhang et
87 al., 2012), aging of BC particles (Zhang et al., 2008), and cloud formation (Ruehl et al.,
88 2016). Dependent of the scientific objectives, the designs of environmental chambers and
89 reactors vary considerably (Zhang et al., 2015). However, few of the previous experimental
90 methods have characterized the evolution of aerosols under the ambient conditions.

91 In this study, we present measurements of aging and hygroscopicity variation of BC
92 particles in Beijing using a quasi-atmospheric aerosol evolution study (QUALITY) chamber
93 (Reed, 2010; Peng et al., 2016). The performance of the QUALITY chamber for mimicking
94 the ambient gaseous concentrations (i.e., the wall loss, and gas mixing rate), ultraviolet
95 transmission, and meteorology parameters (i.e., temperature and relative humidity, RH) has
96 been evaluated.

97 **2 Experimental method**

98 The 1.2 m³ QUALITY chamber was employed to study BC aging under ambient
99 conditions (Fig. 1). The two-layer chamber was comprised of an inner layer of 0.13 mm
100 Perfluoroalkoxy alkane (PFA) Teflon and an outer rigid 5.6 mm thick acrylic shell (Cryo
101 Industries Acrylite, OP-4). Both Acrylite OP-4 Acrylic and PFA Teflon allowed for efficient
102 ultraviolet (UV) transmission in UV-B (280-315 nm) and UV-A (315-400 nm) ranges. When
103 exposed to sunlight, the UV light transmitted through the chamber wall and initiated
104 photochemical reactions inside the chamber.

105 The two individual subdivisions of the QUALITY chamber included a bottom flow
106 chamber, where ambient air was pulled through continuously over each experiment, and an
107 upper reaction chamber, where aging of BC particles occurred (Fig. 1). The two chambers
108 were separated by a 5 μm thick semi-permeable expanded polytetrafluoroethylene (e-PTFE)
109 membrane with high degrees of chemical resistivity, microporosity, nonpolarity, and thermal
110 stability. Gaseous species penetrated the membrane by either bulk gas flow or diffusion. The
111 permeability of the e-PTFE was greater than 90 % for nearly all the tested volatile organic
112 components (Fig. S1) and other gaseous pollutants, i.e., SO₂, NO_x, O₃ and CO. Particles, on
113 the other hand, were blocked from penetration into the reaction chamber. The filtration
114 efficiency of the e-PTFE membrane was measured to be over 99.5% for particles larger than
115 15 nm. During each chamber experiment, ambient air was pulled through the flow chamber
116 continuously and gases in lower chamber permeated through the membrane into the reaction

117 chamber. Hence, an environment that continuously captured the ambient gas concentrations
118 without the presence of ambient particles was created inside the reaction chamber. Since the
119 chamber was continuously exposed to ambient gas concentrations during experiments, gases
120 lost due to reaction, deposition or adsorption to the seed aerosols within the reaction chamber
121 were steadily replenished by the exchange with the flow chamber. Several sampling ports
122 were set at the side of the reaction chamber for injection of seed particles or sampling during
123 experiment.

124 Seed particles were introduced in the reaction chamber via an injection line (Fig. S2). To
125 investigate the growth of BC particles under ambient condition, monodisperse BC particles
126 were injected into the QUALITY chamber. BC particles were generated by incomplete
127 combustion of propane fuel in a custom-made laminar diffusion burner (Santoro et al., 1983;
128 Qiu et al., 2012). The aerosol stream sequentially passed through a 300°C heater to evaporate
129 the semi-volatile organic compounds in the particle phase, a Nafion dryer to remove excess
130 moisture in the flow, and four one-meter-long cylinders containing both alumina spherules
131 coated with potassium permanganate and activated carbon to remove all the gaseous
132 pollutants (i.e., VOCs, H₂S, SO₂, NO_x, O₃). The combined measurement of particle size
133 distribution, density and chemical composition exhibited that organics accounted for less than
134 10% of fresh BC mass concentration. Moreover, the single scattering albedo (SSA) of fresh
135 BC particles was only 0.1, further confirming that few organic coatings existed on fresh BC
136 particles after treatment. The measured removal efficiency in the cylinders for SO₂, NO and
137 NO₂ were 99.2%, 100%, and 99.9%, respectively. The aerosol stream was then introduced
138 through an ionizer and into a differential mobility analyzer (DMA, model 3081, TSI, Inc.)
139 with stable voltage to create a monodisperse BC particles flow. A suite of high time
140 resolution state-of-the-art aerosol instruments simultaneously measured a comprehensive set
141 of BC properties throughout the BC aging process (Table S1). The particle diameter, mass,
142 chemical composition, hygroscopicity and ability as cloud condensation nuclei (CCN) were
143 measured by a scanning mobility particle sizer (SMPS), a differential mobility analyzer–
144 aerosol particle mass analyzer (DMA-APM) system, a high-resolution time-of-flight aerosol
145 mass spectrometer (HR-ToF-AMS), a humidified tandem differential mobility analyzer
146 (HTDMA) system, and a cloud condensation nuclei counter (CCNC), respectively. Detail
147 information of the instruments is provided in the supplementary material and previous
148 publications (DeCarlo et al., 2006; Khalizov et al., 2009). Specific measurement procedures
149 in this study are discussed below.

150 **DMA-APM measurement.** The DMA-APM was used to measure the effective density of
151 BC particles (Pagels et al., 2009). Before any DMA-APM measurement, a SMPS scan was
152 made to obtain the size distribution of particles inside the chamber. The particle size
153 distribution was then fitted with a lognormal Gaussian distribution to derive the peak
154 diameter. During a DMA-APM measurement, the aerosol flow passed through DMA with a
155 fixed voltage to select particles with a fixed diameter. The APM then measured the mass
156 distribution of the selected particles with the same diameter, and the effective density of these
157 particles was obtained by fitting the mass distribution with a normal Gaussian distribution.

158 **DMA-CCN measurement.** Similar to the density measurements, CCN activation was also
159 measured on the basis of monodisperse particles with a peak diameter. A DMA with a fixed

160 voltage selected particles with a peak diameter. Both a CPC and a CCN counter were placed
161 in parallel after the DMA to simultaneously measure the total particle number concentration
162 (N_{cn}) as well as the activated number concentration (N_{ccn}) at a fixed supersaturation. The
163 activation fraction of the BC particles with peak diameter is calculated as:

$$164 \quad f_{CCN} = \frac{N_{CN}}{N_{CCN}} \quad (1)$$

165 Several gradients of supersaturation were set for the CCN counter, with each one being 6
166 min. This method yielded a steeper curve of the CCN activation rate, which was employed to
167 estimate the particle diameter with 50% activation fraction (D_{50}) and the kappa value.

168 Prior to each experiment, the QUALITY chamber was flushed by zero air for more than
169 40 hours to ensure a clean condition and covered with two layers of anti-UV cloth to shield it
170 from sunlight. In the beginning of each experiment, monodisperse BC particles were
171 introduced into the chamber. The injection of BC particles typically lasted for 1 to 2 hours.
172 During the injection period, zero air passed through the bottom chamber continually to
173 remove any possible remaining gaseous pollutant that were removed by the activated carbon.
174 After the injection, ambient air was pulled through the bottom chamber at a flow rate of about
175 50 L min^{-1} for at least half an hour in order to produce a quasi-ambient condition inside the
176 chamber. Finally, the anti-UV cloth was removed, and BC particles underwent aging inside
177 the reaction chamber. A charged zero air stream continuously passed through the space
178 between the two chamber layers to reduce particle wall loss. Various properties of BC
179 particles, including the particle diameter, mass, chemical composition, hygroscopicity, and
180 optical coefficients, were simultaneously measured by a suite of state-of-art aerosol
181 instruments in every 0.5 - 1 hour. Ambient particles and chamber particles were measured
182 alternately every 30 min. The aging experiments lasted for about 2-6 hours depending on the
183 initial BC concentrations and ambient conditions.

184 The BC aging experiments were conducted from August 18th to October 17th, 2013 at an
185 urban site (PKUERS) located on the campus of Peking University in the northwestern Beijing
186 (39.99°N , 116.31°E) (Hu et al., 2012).

187 **3. Characterization and Validation of the QUALITY chamber**

188 **3.1 Wall loss of gases and aerosols**

189 To evaluate the wall loss of both particles and gases in the QUALITY chamber, different
190 gaseous pollutants and particles were introduced into the chamber separately and the decay of
191 their concentration inside the chamber was measured by gas analyzers and SMPS,
192 respectively. All of the ports connected to the ambient air were closed to ensure an enclosed
193 system in the reaction chamber.

194 Particles with different chemical composition exhibited different wall loss rates.
195 Monodispersed BC particles with different diameters showed a small half-lifetime (τ) of
196 about 4-7 hours (Reed, 2010). Aerosol nucleation also occurred inside the chamber likely
197 from organic species (Zhao et al., 2009), which corresponded to a half time of about 3.5
198 hours (Fig. S3), because the nucleated particles inside the chamber was neutral with a slow

199 electrostatic loss to the wall.

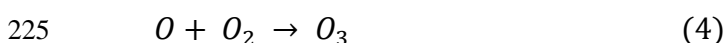
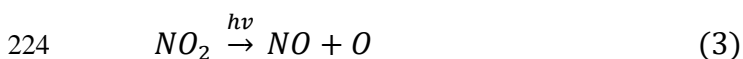
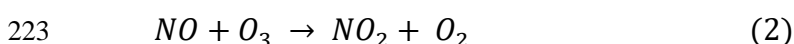
200 Gas species had a longer residence time inside the chamber. Toluene and isoprene did
201 not show obvious wall loss during a two-day experiment (Reed, 2010). O₃, SO₂ and NO_x
202 decreased by 50% inside the chamber after more than 20 hours, suggesting slow loss. Since
203 the loss rate of these primary gaseous pollutants was much slower than the gas exchange rate
204 between the chamber, the loss of gases was replenished by the exchange with the flow
205 chamber.

206 3.2 UV transmission

207 The QUALITY chamber contained two layers of walls, an acrylic shell layer and a PFA
208 Teflon layer. Since the QUALITY chamber used sunlight as the photochemical origin, the
209 transmission spectra of the two-layer walls was of great important for the photochemical
210 reactions inside the reaction chamber.

211 The transmission efficiencies of each material were measured using an ultraviolet-visible
212 (UV-Vis) spectrophotometry (PerkinElmer Inc., model 552). As shown in Figure 2, the
213 Teflon film exhibited stable transmission efficiency of about 60% in the focused wavelength
214 range. The Acrylic shell, however, showed very low transmission efficiency when the
215 wavelength range was shorter than 270 nm, and high transmission efficiency (nearly 90%)
216 when the wavelength range was longer than 300 nm. In general, approximately 60% of the
217 UVA irradiation (315-400 nm range) and 50% of the UVB irradiation (280-315 nm range)
218 penetrated through the chamber walls, allowing photochemical processes to take place in the
219 upper reaction chamber.

220 A NO₂-photolysis experiment was also conducted to characterize the UV transmission of
221 QUALITY chamber. NO₂ was introduced inside the chamber at a clean and sunny day.
222 Reactions among NO₂, NO and O₃ will occur as the following equations:



226 By simultaneously monitoring the concentrations of NO₂, NO and O₃ inside the chamber,
227 the photolysis constant of NO₂, J(NO₂), inside the chamber was estimated. The photolysis
228 constant of NO₂ inside the chamber was on average 55% of that in the ambient air, in
229 agreement with the transmission spectra measurement, further confirming that the two-layer
230 chamber walls allowed 50% - 60% of solar irradiation in the UV range into the reaction
231 chamber.

232 3.3 Gas concentrations in the reaction chamber

233 To investigate the VOCs concentration inside the reaction chamber after the injection of
234 BC particles, both the chamber and ambient air were sampled with VOC canisters just before
235 the BC aging experiment started. These canisters were then analyzed by a gas
236 chromatography-mass spectrometer / flame ionization detector (GC-MS/FID, HP inc.) system

237 (Liu et al., 2008). The concentrations of VOCs containing 4 or more carbons are illustrated in
238 Figure 3. Slightly higher concentrations of several VOCs in the QUALITY chamber, e.g.,
239 n-butane, n-pentane, toluene, were observed compared with those in the ambient air, due to
240 the co-injection of a small amount of VOCs together with BC particles into the chamber.
241 Nevertheless, the average increase of the VOC concentrations was only 16% or 0.1 ppb for
242 all focused VOCs species, with the largest increase of 35% or 0.36 ppb, suggesting the
243 insignificant influence of soot burner on VOCs concentrations and SOA formation in the
244 chamber.

245 Additional experiments were conducted to characterize the exchange time scale for gases
246 in the QUALITY chamber. The chamber was firstly cleaned and flushed with zero air for 40
247 hours. Ambient air was then pulled through the bottom chamber (Fig. 4), and the
248 concentrations of gaseous pollutants, including O₃, NO_x, CO, and SO₂, were measured
249 alternatively in the ambient air and in the upper chamber. In the beginning of the experiment
250 the concentrations of all gaseous species were lower than those in the ambient air. After
251 ambient air was pulled through the bottom chamber (labeled as the black dots in Fig. 4), the
252 concentrations of the gaseous pollutants in the reaction chamber increased sharply. For
253 example, the CO concentration inside the chamber was approximate 70% of that the ambient
254 concentration after 30-min mixing. The concentrations inside the chamber and ambient air
255 exhibited little difference after 1-hour mixing (Fig. 4c). Hence, the QUALITY chamber well
256 replicated the ambient gas concentrations. The gas exchange rate between the bottom and
257 reaction chambers was calculated to be approximately 0.06 min⁻¹.

258 **3.4 Temperature and RH**

259 The greenhouse effect for an outdoor chamber typically increases the temperature and
260 decreases the RH inside the chamber. For the QUALITY chamber, however, heat produced
261 by the greenhouse effect inside the chamber was effectively taken away as the ambient air
262 continuously passed through the bottom chamber and exchanged with air in the upper
263 chamber. As illustrated in Figure 5, there was little difference in temperature or RH inside
264 and outside the chamber, when the chamber experiments lasted for more than 1 hour,
265 suggesting that the QUALITY chamber effectively captured the ambient temperature and
266 RH.

267 **3.5 Sulfuric acid production**

268 Though gaseous pollutants such as SO₂, O₃, NO_x and VOCs penetrated into the reaction
269 chamber from the bottom chamber through the semi-permeable membrane, low volatile and
270 sticky gases, i.e., sulfuric acid, were unlikely to penetrate through the membrane (Fortner et
271 al., 2004). To characterize sulfuric acid production inside the chamber, we conducted a
272 special experiment by pulling ambient air into the bottom chamber, while injecting SO₂
273 directly into the reaction chamber. The experiment was conducted around noon on a clean
274 day when the O₃ concentration was around 50 ppb. A custom-made atmospheric pressure-ion
275 drift chemical ionization mass spectrometry (AP-ID-CIMS) (Fortner et al., 2004; Zheng et al.,
276 2010) was used to directly measure the concentration of gaseous sulfuric acid. Fig. 6 shows a
277 good correlation between SO₂ and gaseous H₂SO₄ inside the chamber, suggesting that the
278 QUALITY chamber well simulated the formation of low volatile gaseous species and hence

279 the photochemical processes.

280 4. BC evolution in the QUALITY chamber

281 Time series of the ambient PM, gas, and meteorology parameters of ambient air during
282 chamber experiment period are illustrated in Figure 7. Except for the last experiment
283 conducted on October 17th, all experiments were conducted between August 18th, 2013 and
284 September 21st, 2013. During this period in Beijing, strong photochemical reactions and
285 frequent heavy pollution events occurred (Huang et al., 2010; Zhao et al., 2013). Temperature
286 and RH during this period (August 18th to September 21st) varied from 24°C to 38°C and
287 from 20 % to 90 %, respectively. The average concentration of PM_{2.5}, SO₂ and NO_x were 60
288 ± 45 µg m⁻³, 3.2 ± 2.6 ppb and 33.9 ± 20.9 ppb, respectively.

289 The red shaded areas in Figure 7 represent the period of the nine chamber experiments.
290 In general, chamber experiments were conducted in the afternoon of relatively clean and
291 sunny days, when strong solar radiation led to fast photochemical reactions. There were two
292 experiments conducted under polluted days, e.g., the experiments on October 22nd and
293 September 11th. Table 1 summaries the conditions of the experiments. Totally, 10 BC aging
294 experiments were conducted, including four experiments using BC particles with initial
295 mobility diameter of 100 nm, three experiments using 150 nm particles, and three using 220
296 nm BC particles. The average concentrations of PM_{2.5} and NO_x over each chamber
297 experiment were only 9 to 69 µg m⁻³ and 9 to 41 ppb, respectively. The concentrations of
298 VOCs, such as toluene and m/p-xylene were relatively low during most of the experiments
299 compared with severe pollution episodes in Beijing (Guo et al., 2014). *J*(O¹D) and O₃
300 exhibited higher values during the chamber experiments. The average *J*(O¹D) values and O₃
301 concentrations ranged from 3.2 × 10⁻⁶ s⁻¹ to 21.1 × 10⁻⁶ s⁻¹ and 26 to 92 ppb, respectively.

302 4.1 BC growth

303 To quantify the growth of BC particles, several parameters were used to describe the
304 properties of BC particle, including the effective density, mobility diameter and mass
305 equivalent diameter. Material density (ρ_m) is the average density of the solid and liquid
306 material in the particle. Assuming that the volume of the species coexisting in an individual
307 particle does not change upon mixing, the density of an internally mixed particle is calculated
308 using the material densities and mass concentrations of particulate constituents (DeCarlo et
309 al., 2004; Pagels et al., 2009),

$$310 \quad \rho_m = \frac{\sum_{species} m_i}{\sum_{species} V_i} = \frac{\sum_{species} m_i}{\sum_{species} \frac{m_i}{\rho_i}} = \frac{\sum_{species} MC_i}{\sum_{species} \frac{MC_i}{\rho_i}} \quad (5)$$

311 where ρ_i is the material density of species i , V_i is its volume, m_i is its mass, and MC_i is its
312 mass concentration. This approach is based on the assumption that there is no void space
313 enclosed within the particle envelope. Hence, the material density is larger than the true
314 particle density with internal voids in particles. In this study, ρ_m is calculated from the
315 chemical composition of coating materials measured by AMS. A value of 1.35 for the
316 material density of SOA formed during chamber experiment was obtained by directly
317 measuring the density of newly form particles inside the chamber via DMA-APM system.

318 Effective density (ρ_{eff}) is defined as the ratio of the measured particle mass (m_p) to the
 319 particle volume calculated assuming a spherical particle with a diameter equal to the
 320 measured mobility diameter (D_m) (DeCarlo et al., 2004; Pagels et al., 2009; Xue et al.,
 321 2009b):

$$322 \quad \rho_{eff} = \frac{6m_p}{\pi D_m^3} \quad (6)$$

323 In this study, m_p of BC particles was measured by the APM and mobility diameter (D_m)
 324 was measured by the DMA. The effective density reflects the information on both particle
 325 density and shape. If particles are spherical in the absence of internal void, the effective
 326 density equals the material density. If particles are non-spherical, the calculated volume and
 327 volume concentration are larger than the true values, and the effective density is less than true
 328 particle and the material density.

329 Figure 8 exhibits the change of particle density and diameter in three typical BC aging
 330 experiments using BC particles with the initial diameter of 100 nm, 150 nm, and 220 nm
 331 (experiment #4, #5, and # 9 in Table 1, respectively). The average PM_{2.5} concentrations in
 332 these experiments were 40, 27, and 12 $\mu\text{g m}^{-3}$, respectively, suggesting relatively clean
 333 conditions during the experiments. In all three experiments, aging of BC particles occurred
 334 between 13:00-14:00 and 17:00 in the afternoon. The highest $J(\text{O}^1\text{D})$ value varied from 1.7 to
 335 $2.4 \times 10^{-5} \text{ s}^{-1}$ and decreased generally over the experiment period. Average O₃ concentrations
 336 during the three experiments were 68, 83 and 54 ppb, respectively, indicating strong
 337 oxidation during the experiment periods.

338 For three initial D_m of 100 nm, 150 nm, and 200 nm, the ranges of effective density of
 339 fresh BC particles in each experiment were 0.43-0.50 g cm^{-3} , 0.34-0.34 g cm^{-3} , and 0.24-0.32
 340 g cm^{-3} , respectively, indicating highly fractal BC aggregates (Zhang et al., 2008). The small
 341 variation for particles with the same D_m also demonstrates the consistency of fresh BC
 342 morphology in different experiments. After aging of 1 hour, ρ_{eff} of BC particles in the three
 343 experiments increased to approximately 1.2 g cm^{-3} , suggesting that formation of the
 344 secondary components changed the morphology from chain-like BC particles into a more
 345 compact shape (Zhang et al., 2008; Peng et al., 2016). The morphology change was further
 346 confirmed by a decrease of D_m , particularly for larger BC particles that were more fractal
 347 than smaller BC particles.

348 As the D_m is largely influenced by the particle morphology, we utilized the parameter of
 349 mass equivalent diameter (D_{me}) to describe the growth of particles. Based on the mobility
 350 diameter (D_m), material density (ρ_m) and effective density (ρ_{eff}), the D_{me} is calculated
 351 assuming that particles are compact and with a spherical morphology (DeCarlo et al., 2004):

$$352 \quad D_{me} = \sqrt[3]{\frac{\rho_{eff}}{\rho_m}} D_m \quad (7)$$

353 The change in the mass equivalent diameter (ΔD_{me}) during BC aging is defined as the total
 354 coating thickness, and the ratio of the total coating thickness to the initial mass equivalent
 355 diameter ($\Delta D_{me}/D_{me,0}$) is defined as the coating fraction.

356 The initial D_{me} of fresh BC particles with initial D_m of 100, 150 and 220 nm were 61, 84
357 and 114 nm, respectively. In contrast to the mobility diameter, D_{me} increased continuously
358 during the entire experiment. After 3-4 hours, D_{me} in the three experiments increased to 133,
359 169 and 197 nm, respectively (Fig. 8), with the average growth rates of 19, 29 and 31 nm h⁻¹.
360 Higher growth rates in D_{me} occurred around noontime, when the $J(O^1D)$ value was higher and
361 the photochemical reaction was stronger. On the other hand, much less growth rate was
362 observed during late afternoon or with cloud coverage (As shown in Fig. 8c at 15:00),
363 indicating that the growth was driven by photochemical reactions.

364 The increases of the particle density and diameter in all the experiments are summarized
365 in Table 2. Fast aging of BC particles occurred in all experiments. The total growth of D_{me}
366 ranged from 40 nm to 152 nm within 3-6 hours, with an average growth of 73 nm. The
367 average growth rate was 26 ± 11 nm h⁻¹, demonstrating large secondary production under the
368 ambient conditions in Beijing. The largest growth rate ($\Delta D_{me} = 152$ nm) was observed in
369 experiment #8, when solar irradiation was the strongest among all experiments (Table 2).

370 Correlation analysis was made between the average growth rate of BC particles ($\Delta D_{me}/\Delta t$)
371 with O_3 , $PM_{2.5}$, $J(O^1D)$, and temperature during the different experiments (Fig. S4). The
372 growth rate of BC particles exhibits no correlation with O_3 concentration ($R^2=0.00$), weak
373 negative correlation with $PM_{2.5}$ concentration ($R^2 = 0.25$), and strong positive correlations
374 with $J(O^1D)$ ($R^2 = 0.80$) and temperature ($R^2 = 0.67$), indicating the importance of
375 photochemical production on the BC coating materials.

376 4.2 Chemical composition of coating materials

377 Particle composition measurements by AMS during chamber experiments reveal a
378 majority of coating materials (above 90%) as SOA (Fig. 8). The concentration of SOA inside
379 the chamber reached up to $9 \mu\text{g m}^{-3}$ in several experiments, suggesting fast formation of SOA
380 via gas phase oxidation of VOCs. The SOA formation in Beijing is likely attributed to a large
381 amount of anthropogenic aromatic VOCs (Peng et al., 2017).

382 The elemental compositions of OA inside the chamber, i.e., the oxygen to carbon (O/C)
383 ratio and the hydrogen to carbon (H/C) ratio, were calculated based on the updated ambient
384 calibrations (He et al., 2011; Canagaratna et al., 2015). The H/C and O/C ratios of organics
385 for coating on BC particles exhibit notable trends during the aging process. Figure 9A shows
386 an example of the evolution of H/C and O/C ratios in experiment #8. The data were corrected
387 for the CO_2 concentration in the chamber, which were introduced into the chamber with BC
388 particles and influenced the abundance of $m/z = 28$ and 44 in the AMS mass spectra. The H/C
389 ratio decreased from 1.73 to 1.45 over six hours. Accordingly, the O/C ratio increased from
390 0.32 to 0.50 during the same time, revealing that further oxidation of SOA occurred in the
391 latter part of the experiment. The lower final O/C ratio in the chamber experiment (0.5) than
392 that under the ambient conditions (Hu et al., 2016) implies that there is oxidation on a longer
393 timescale or by the aqueous pathway for the formation of highly oxidized SOA in the ambient
394 air (Zhang et al., 2015).

395 Furthermore, the mass spectra of OA inside the chamber shows strong correlation with
396 the less-oxidized oxygenated organic aerosols (LO-OOA) derived from field measurements

397 in Beijing (Hu et al., 2016), which likely arose from oxidation of aromatic VOCs emitted
398 from vehicles (Peng et al., 2017). The correlation coefficient (R^2) initially was 0.88 and raised
399 to 0.99 sharply (Fig. 9B), indicating that the chamber well simulated the formation of
400 LO-OOA.

401 In our study, the secondary inorganic aerosols, i.e., sulfate, nitrate and ammonium, only
402 accounted for less than 10% of the coating materials on BC particles. This is consistent with
403 the previous studies showing that the concentration of organics is much larger than those of
404 sulfate and nitrate during the early stage of haze development in Beijing (Guo et al., 2014).
405 The low observed sulfate concentration in this study suggests that the gas phase formation of
406 sulfuric acid was unimportant under our experimental conditions. On the other hand, it has
407 been shown that the aqueous-phase reactions represent the dominate pathway for sulfate
408 formation in Beijing (Guo et al., 2010; Wang et al., 2016).

409 NO_2 has a higher reaction coefficient with the OH radical ($8 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$)
410 than SO_2 (Zhang et al., 2015). Nitrate acid formed in the gas phase is transformed into nitrate
411 salts by the reaction with ammonia in the equilibrium process:



413 The equilibrium of this reaction is highly depended on ambient temperature and RH (Zheng
414 et al., 2008). In this study, chamber experiments were conducted in the afternoon with high
415 temperature and low RH (Table 1), which shifted the thermodynamic equilibrium to the gas
416 phase.

417 **4.3 Hygroscopicity evolution**

418 **HTDMA measurement**

419 The hygroscopic growth factors (HGF) of particles in each experiment were
420 continuously measured by the HTDMA system and corrected for the reference “dry”
421 diameters,

$$422 \quad \text{HGF} = \frac{D_{wet,t}/D_{m,t}}{D_{dry,0}/D_{m,0}} \quad (9)$$

423 where D_m is the mobility diameter of fresh or coated particles at dry condition, D_{dry} is the
424 mobility diameter of particles after experiencing a low humidity (below 30%) cycle in
425 HTDMA, and D_{wet} is the mobility diameter of particles after experiencing a high humidity
426 cycle (87%) in HTDMA.

427 Figure 10 shows the hygroscopicity variation of BC particles with the initial mobility
428 diameter ($D_{m,0}$) of 100 nm and 150 nm. The measured HGF of 0.999 - 1.004 for fresh BC
429 particles suggests high hydrophobicity, consistent with the previous studies (Khalizov et al.,
430 2009; Weingartner et al., 1997). After exposed to sunlight and ambient gaseous pollutants for
431 several hours, the HGF of these BC particles increased to 1.02-1.08 at the end of each
432 experiment. The HGF value varied with the total growth (ΔD_{me}) of BC particles, but was
433 constant at the same ΔD_{me} for different experiments (Fig. 10). The final HGF values shown in

434 Figure 10 (1.02-1.08) were much lower than those in previous laboratory studies (Khalizov et
435 al., 2009; Tritscher et al., 2011) but similar to the low hygroscopic fraction in field
436 observations (Swietlicki et al., 2008), even for growth of particle size up to 90 nm in our
437 experiments.

438 The HGF is affected by many factors, e.g., the particle chemical composition and
439 morphology as well as RH (Qiu et al., 2012). The hygroscopicity of BC particles coated with
440 inorganic components, i.e., sulfuric acid (Khalizov et al., 2009), is significantly higher than
441 that coated by organic compounds (Tritscher et al., 2011). In this study, the major component
442 of the coating substance was LO-OOA with a O/C ratio about 0.5. The low oxygen content of
443 SOA coated on BC particles explains the low hygroscopicity (Jimenez et al., 2009),
444 indicating that coating of BC particles during the early stage haze development in Beijing
445 does not considerably increase the particle hygroscopicity.

446 The morphology of BC particles directly affects the HGF. As illustrated in Figure 11,
447 when the ΔD_{me} was 18 nm and 22 nm for 100 nm and 150 nm BC particles, respectively, the
448 HGF decreased slightly to about 0.99, suggesting that a thin layer of coatings on BC particles
449 decreased the particle diameter, even though a certain amount of water absorbed by BC
450 particles increased the particle mass. The surface tension of the water layer produced an
451 inward force on the “chain-like” branches of BC particles, leading to particle reconstruction,
452 and a more compact morphology. Such change was also identified in laboratory studies
453 (Weingartner et al., 1997; Tritscher et al., 2011; Qiu et al., 2012). In this study, the BC
454 particles became spherical when ΔD_{me} was 30 nm and 40 nm for particles with initial D_m of
455 100 nm and 150 nm, respectively (Peng et al., 2016). Therefore, when ΔD_{me} was large, the
456 HGF value was not influenced by reconstruction.

457 CCN measurements and κ closure

458 The CCN activation fraction (f_{CCN}) of BC particles at different supersaturation during two
459 typical experiments is illustrated in Figure 11. Fresh BC particles were not activated even at
460 very high supersaturation conditions (0.7%). With aging, f_{CCN} rapidly raised to nearly 100%
461 at high supersaturation (0.7% for experiment #4 and 0.6% for experiment #6). After several
462 hours, BC particles became CCN at lower supersaturation. The f_{CCN} at 0.4 supersaturation
463 (Fig. 11a in experiment #4) and 0.3 supersaturation (Fig. 11b in experiment #6) exceeded 50%
464 before the end of these two experiments, suggesting that aging increases the ability of BC
465 particles to become CCN (Wittbom et al., 2014) and a large amount of coatings results in
466 activation at lower supersaturation.

467 To further investigate the hygroscopicity of BC particles and combine the measurements
468 using HTDMA and CCN, we evaluated the hygroscopicity parameter, kappa (κ) (Petters and
469 Kreidenweis, 2007). The approximate relationship between the dry particle mass equivalent
470 diameter (D_{me}), the critical saturation ratio (S_c) and the apparent κ value of particles is
471 describe as:

$$472 \quad \kappa = \frac{4A^3}{27D_{me}\ln^2 S_c} \quad (10)$$

473 where A is a parameter that includes several features of the solvent,

474
$$A = \frac{4\sigma_{s/a}M_w}{RT\rho_w} \quad (11)$$

475 M_w is the molecular weight of water, ρ_w is the density of water, $\sigma_{s/a}$ is the surface tension of
 476 the solution/air interface, R is the universal gas constant, and T is temperature.

477 In addition to the supersaturated condition, the κ theory also adopts the form for the
 478 subsaturated condition, using the HGF from HTDMA measurement and RH:

479
$$\frac{RH}{\exp\left(\frac{A}{D_{wet}}\right)} = \frac{HGF^3 - 1}{HGF^3 - (1 - \kappa)} \quad (12)$$

480 where D_{wet} is the wet diameter of particles.

481 The apparent κ values of BC particles calculated by HTDMA (κ_{HTDMA}) and CCN
 482 (κ_{CCN}) are shown in Figure 11. The κ of fresh BC particles was near zero. With aging,
 483 SOA coated on BC particles increased the κ_{HTDMA} and κ_{CCN} to approximately 0.04,
 484 although the κ_{HTDMA} and κ_{CCN} exhibited difference features. A slightly higher κ_{CCN}
 485 than κ_{HTDMA} at the beginning of aging was identified, attributed to reconstruction of BC
 486 particles after humidified and underestimation of HGF and thus the κ_{HTDMA} value. Such a
 487 difference between κ_{CCN} and κ_{HTDMA} was also observed in previous studies (Tritscher
 488 et al., 2011; Martin et al., 2013). Nevertheless, the apparent κ values from two both methods
 489 were comparable at the end of both experiments.

490 Assuming that a simple mixing rule is applicable to coated BC particles, the κ for
 491 coating materials can be calculated based on the volume fraction of BC and SOA:

492
$$\kappa = \sum_i \varepsilon_i \kappa_i \quad (13)$$

493 where ε_i represents the volume fraction of species i . Since the SOA formed inside the
 494 chamber was not highly hygroscopic, some of the SOA components might not be able to
 495 solve in water droplets (Petters and Kreidenweis, 2008), leading to the underestimation of the
 496 κ values of the coating materials in this study.

497 The κ values of the coating materials were 0.04 at the end of our experiments for both
 498 CCN and HTDMA method, much lower than that of ambient aerosols in Beijing (Gunthe et
 499 al., 2011; Wu et al., 2016) and those of SOA in previous chamber studies (Jimenez et al.,
 500 2009; Tritscher et al., 2011; Martin et al., 2013). As discussed above, the coating substances
 501 on BC particles were mainly SOA formed from photochemical oxidation. The κ of SOA
 502 depends on the oxidation degree, which is correlated to the O/C ratio (Jimenez et al., 2009;
 503 Massoli et al. 2010). The O/C ratio of the coating SOA was 0.5 in our experiment,
 504 corresponding to the O/C ratio of approximately 0.4 in Jimenez et al. (2009) and Massoli et al.
 505 (2010) duo the utilize of updated AMS calibration method in this study. The κ value of
 506 coating materials here is in general slightly lower than those of SOA with similar oxidation
 507 degree (O/C ratio) formed from different VOCs precursors (Jimenez et al., 2009; Massoli et
 508 al. 2010). As motion above, the κ values of the coating materials might be underestimated
 509 duo to the ignore of solubility of the SOA products. Also, our experiments represent the aging

510 of BC under typical urban condition, where variety of anthropogenic VOCs, i.e., toluene,
511 xylene, TMB and small molecular PAHs, contribute significantly to the SOA formation.
512 Some products of these VOC precursors may exhibit low hygroscopicity.

513 **5 Conclusions**

514 In this paper, we present measurements of aging and hygroscopicity of BC particles in
515 Beijing using the QUALITY chamber. The unique two sub-chamber design facilitates the
516 evaluation of aging of BC particles under ambient conditions, by mimicking the ambient
517 gaseous concentrations without the presence of ambient aerosols. High UV transmission
518 efficiency (50-60%) and negligible wall loss of primary gaseous pollutants are shown for the
519 chamber performance. The validation experiments demonstrate little differences in the
520 primary gas concentrations, temperature, and RH between the chamber and the atmosphere,
521 suggesting that the chamber captures the evolution of ambient conditions. In addition, our
522 results show sulfuric acid production correlated with SO₂, indicating that the chamber well
523 simulates photochemical-driven formation of low volatile gaseous species by the hydroxyl
524 radical.

525 BC aging experiments were performed using the QUALITY chamber in Beijing. Fast
526 growth of BC particles (on average 26 ± 11 nm h⁻¹) was observed, and SOA was identified as
527 the dominate component of the coating materials on BC particles, while inorganic species,
528 such as sulfate and nitrate, were unimportant under our experimental condition and timescale.

529 The HGF of BC particles exhibited a very low value (1.02-1.08) after several hours aging.
530 A slight decrease of HGF with a thin coating layer indicated reconstruction of BC particles
531 after humidified. Also, a very low kappa value (0.035) for BC particles at both subsaturation
532 and supersaturation conditions were found, with HTDMA and CCN measurements. Hence,
533 our results indicate that initial photochemical aging of BC particles does not appreciably alter
534 the particle hygroscopicity in Beijing.

535

536 **Acknowledgement**

537 This work was supported by National Natural Science Foundation of China (91544214,
538 41421064), the National Basic Research Program, China Ministry of Science and Technology
539 (Grant 2013CB228503, Grant 2013CB955801), National Natural Science Foundation of
540 China (Grant 21190052), and the China Ministry of Environmental Protection's Special
541 Funds for Scientific Research on Public Welfare (Grant 20130916). R.Z. acknowledged
542 support from the Robert A. Welch Foundation (Grant A-1417) and Houston Advanced
543 Research Center. We thanked Wei Hu and Zhaoheng Gong for their assistance with the AMS
544 data analysis, Wentai Chen and Yue Li for providing VOCs data.

545

546

547 **References**

- 548 Bond, T. C., Doherty, S. J., Fahey, D. W., Forster, P. M., Berntsen, T., DeAngelo, B. J.,
549 Flanner, M. G., Ghan, S., Kärcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P. K.,
550 Sarofim, M. C., Schultz, M. G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S.,
551 Bellouin, N., Guttikunda, S. K., Hopke, P. K., Jacobson, M. Z., Kaiser, J. W., Klimont, Z.,
552 Lohmann, U., Schwarz, J. P., Shindell, D., Storelvmo, T., Warren, S. G., and Zender, C. S.:
553 Bounding the role of black carbon in the climate system: A scientific assessment, *Journal*
554 *of Geophysical Research: Atmospheres*, 118, 5380-5552, 10.1002/jgrd.50171, 2013.
- 555 Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P.,
556 Hildebrandt Ruiz, L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue,
557 N. M., Jayne, J. T., and Worsnop, D. R.: Elemental ratio measurements of organic
558 compounds using aerosol mass spectrometry: characterization, improved calibration, and
559 implications, *Atmos. Chem. Phys.*, 15, 253-272, doi:10.5194/acp-15-253-2015, 2015.
- 560 Claeys, M.: Formation of Secondary Organic Aerosols Through Photooxidation of Isoprene,
561 *Science*, 303, 1173-1176, doi:10.1126/science.1092805, 2004.
- 562 DeCarlo, P. F., Slowik, J. G., Worsnop, D. R., Davidovits, P., and Jimenez, J. L.: Particle
563 Morphology and Density Characterization by Combined Mobility and Aerodynamic
564 Diameter Measurements. Part 1: Theory, *Aerosol Sci. Tech.*, 38, 1185-1205,
565 10.1080/027868290903907, 2004.
- 566 DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C.,
567 Gonin, M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez, J. L.:
568 Field-deployable, high-resolution, time-of-flight aerosol mass spectrometer, *Anal. Chem.*,
569 78, 8281-8289, doi:10.1021/ac061249n, 2006.
- 570 DeCarlo, P. F., Ulbrich, I. M., Crounse, J., de Foy, B., Dunlea, E. J., Aiken, A. C., Knapp, D.,
571 Weinheimer, A. J., Campos, T., Wennberg, P. O., and Jimenez, J. L.: Investigation of the
572 sources and processing of organic aerosol over the Central Mexican Plateau from aircraft
573 measurements during MILAGRO, *Atmos. Chem. Phys.*, 10, 5257-5280,
574 doi:10.5194/acp-10-5257-2010, 2010.
- 575 Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud
576 droplets and aqueous particles (aqSOA): a review of laboratory, field and model studies,
577 *Atmos. Chem. Phys.*, 11, 11069-11102, doi:10.5194/acp-11-11069-2011, 2011.
- 578 Fortner, E. C., Zhao, J., and Zhang, R. Y.: Development of ion drift-chemical ionization mass
579 spectrometry, *Anal. Chem.*, 76, 5436-5440, doi:10.1021/ac0493222, 2004.
- 580 Gunthe, S. S., Rose, D., Su, H., Garland, R. M., Achtert, P., Nowak, A., Wiedensohler, A.,
581 Kuwata, M., Takegawa, N., Kondo, Y., Hu, M., Shao, M., Zhu, T., Andreae, M. O., and
582 Pöschl, U.: Cloud condensation nuclei (CCN) from fresh and aged air pollution in the
583 megacity region of Beijing, *Atmos. Chem. Phys.*, 11, 11023-11039,
584 doi:10.5194/acp-11-11023-2011, 2011.
- 585 Guo, S., Hu, M., Lin, Y., Gomez-Hernandez, M., Zamora, M. L., Peng, J. F., Collins, D. R.,
586 and Zhang, R. Y.: OH-Initiated Oxidation of m-Xylene on Black Carbon Aging, *Environ.*
587 *Sci. Technol.*, 50, 8605-8612, doi:10.1021/acs.est.6b01272, 2016.
- 588 Guo, S., Hu, M., Wang, Z. B., Slanina, J., and Zhao, Y. L.: Size-resolved aerosol
589 water-soluble ionic compositions in the summer of Beijing: implication of regional
590 secondary formation, *Atmos. Chem. Phys.*, 10, 947-959, 2010.

591 Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M.,
592 Zeng, L., Molina, M. J., and Zhang, R.: Elucidating severe urban haze formation in China,
593 *P. Natl. Acad. Sci. USA*, 111, 17373-17378, doi:10.1073/pnas.1419604111, 2014.

594 He, L.-Y., X.-F. Huang, L. Xue, M. Hu, Y. Lin, J. Zheng, R. Zhang, and Y.-H. Zhang:
595 Submicron aerosol analysis and organic source apportionment in an urban atmosphere in
596 Pearl River Delta of China using high-resolution aerosol mass spectrometry, *J. Geophys.*
597 *Res.* 116, D12304, doi:10.1029/2010JD014566, 2011.

598 Hu, M., Peng, J., Sun, K., Yue, D., Guo, S., Wiedensohler, A., and Wu, Z.: Estimation of
599 Size-Resolved Ambient Particle Density Based on the Measurement of Aerosol Number,
600 Mass, and Chemical Size Distributions in the Winter in Beijing, *Environ. Sci. Technol.*,
601 46, 9941-9947, doi:10.1021/es204073t, 2012.

602 Hu, W., Hu, M., Hu, W., Jimenez, J. L., Yuan, B., Chen, W., Wang, M., Wu, Y., Chen, C.,
603 Wang, Z., Peng, J., Zeng, L., and Shao, M.: Chemical composition, sources and aging
604 process of sub-micron aerosols in Beijing: contrast between summer and winter, *J.*
605 *Geophys. Res.- Atmos.*, 121, 1955-1977 doi:10.1002/2015jd024020, 2016.

606 Huang, X. F., He, L. Y., Hu, M., Canagaratna, M. R., Sun, Y., Zhang, Q., Zhu, T., Xue, L.,
607 Zeng, L. W., Liu, X. G., Zhang, Y. H., Jayne, J. T., Ng, N. L., and Worsnop, D. R.: Highly
608 time-resolved chemical characterization of atmospheric submicron particles during 2008
609 Beijing Olympic Games using an Aerodyne High-Resolution Aerosol Mass Spectrometer,
610 *Atmos. Chem. Phys.*, 10, 8933-8945, doi:10.5194/acp-10-8933-2010, 2010.

611 Intergovernmental Panel on Climate Change (IPCC), *Climate Change 2013: The Physical*
612 *Science Basis*, Cambridge Univ. Press, Cambridge, UK, 2013.

613 Jacobson, M. Z.: Strong radiative heating due to the mixing state of black carbon in
614 atmospheric aerosols, *Nature*, 409, 695-697, doi:10.1038/35055518, 2001.

615 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H.,
616 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I.
617 M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A.,
618 Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J.,
619 Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J.,
620 Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K.,
621 Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo,
622 D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J.
623 Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C.,
624 Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E.,
625 Baltensperger, U., and Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere,
626 *Science*, 326, 1525-1529, doi:10.1126/science.1180353, 2009.

627 Kalberer, M.: Identification of Polymers as Major Components of Atmospheric Organic
628 Aerosols, *Science*, 303, 1659-1662, doi:10.1126/science.1092185, 2004.

629 Khalizov, A. F., Xue, H., and Zhang, R.: Enhanced light absorption and scattering by carbon
630 soot aerosols internally mixed with sulfuric acid, *J. Phys. Chem.*, 113, 1066-1074,
631 doi:10.1021/jp807531n, 2009.

632 Khalizov, A. F., Zhang, R., Zhang, D., Xue, H., Pagels, J., and McMurry, P. H.: Formation of
633 highly hygroscopic soot aerosols upon internal mixing with sulfuric acid vapor, *J.*
634 *Geophys. Res.*, 114, doi:10.1029/2008jd010595, 2009.

635 Khalizov, A. F., Lin, Y., Qiu, C., Guo, S., Collins, D., and Zhang, R.: Role of OH-Initiated
636 Oxidation of Isoprene in Aging of Combustion Soot, *Environ. Sci. Technol.*, 47,
637 2254-2263, doi:10.1021/es3045339, 2013.

638 Kuwata, M., Kondo, Y., Mochida, M., Takegawa, N., and Kawamura, K.: Dependence of
639 CCN activity of less volatile particles on the amount of coating observed in Tokyo, *J.*
640 *Geophys. Res.*, 112, doi:10.1029/2006jd007758, 2007.

641 Liu, S., Aiken, A. C., Gorkowski, K., Dubey, M. K., Cappa, C. D., Williams, L. R., Herndon,
642 S. C., Massoli, P., Fortner, E. C., Chhabra, P. S., Brooks, W. A., Onasch, T. B., Jayne, J. T.,
643 Worsnop, D. R., China, S., Sharma, N., Mazzoleni, C., Xu, L., Ng, N. L., Liu, D., Allan, J.
644 D., Lee, J. D., Fleming, Z. L., Mohr, C., Zotter, P., Szidat, S., and Prevot, A. S. H.:
645 Enhanced light absorption by mixed source black and brown carbon particles in UK
646 winter, *Nat. Commun.*, 6, Artn 8435, doi:10.1038/Ncomms9435, 2015.

647 Liu, Y., Shao, M., Fu, L. L., Lu, S. H., Zeng, L. M., and Tang, D. G.: Source profiles of
648 volatile organic compounds (VOCs) measured in China: Part I, *Atmospheric Environment*,
649 42, 6247-6260, doi:10.1016/j.atmosenv.2008.01.070, 2008.

650 Ma, Y., Brooks, S. D., Vidaurre, G., Khalizov, A. F., Wang, L., and Zhang, R.: Rapid
651 modification of cloud-nucleating ability of aerosols by biogenic emissions, *Geophys. Res.*
652 *Lett.*, 40, 6293-6297, doi:10.1002/2013gl057895, 2013.

653 Martin, M., Tritscher, T., Jurányi, Z., Heringa, M. F., Sierau, B., Weingartner, E., Chirico, R.,
654 Gysel, M., Prévôt, A. S. H., Baltensperger, U., and Lohmann, U.: Hygroscopic properties
655 of fresh and aged wood burning particles, *J. Aerosol Sci.*, 56, 15-29,
656 doi:10.1016/j.jaerosci.2012.08.006, 2013.

657 Massoli, P., Lambe, A. T., Ahern, A. T., Williams, L. R., Ehn, M., Mikkilä, J., Canagaratna, M.
658 R., Brune, W. H., Onasch, T. B., Jayne, J. T., Petäjä, T., Kulmala, M., Laaksonen, A., Kolb,
659 C. E., Davidovits, P., Worsnop, D. R.: Relationship between aerosol oxidation level and
660 hygroscopic properties of laboratory generated secondary organic aerosol (SOA) particles.
661 *Geophys. Res. Lett.*, 37, L24801, doi:10.1029/2010GL045258, 2010. Moffet, R. C., and
662 Prather, K. A.: In-situ measurements of the mixing state and optical properties of soot
663 with implications for radiative forcing estimates, *P. Natl. Acad. Sci. USA*, 106,
664 11872-11877, doi:10.1073/pnas.0900040106, 2009.

665 Pagels, J., McMurry, P.H., Khalizov, A.F., and Zhang, R.: Processing of soot by controlled
666 sulphuric acid and water condensation—Mass and mobility relationship, *Aerosol Sci.*
667 *Tech.* 43, 629–640, 2009.

668 Peng, J., Hu, M., Du, Z., Wang, Y., Zheng, J., Zhang, W., Yang, Y., Qin, Y., Zheng, R., Xiao,
669 Y., Wu, Y., Lu, S., Wu, Z., Guo, S., Mao, H., and Shuai, S.: Gasoline aromatic: a critical
670 determinant of urban secondary organic aerosol formation, *Atmos. Chem. Phys. Discuss.*,
671 doi:10.5194/acp-2017-254, 2017.

672 Peng, J. F., Hu, M., Guo, S., Du, Z. F., Zheng, J., Shang, D. J., Zamora, M. L., Zeng, L. M.,
673 Shao, M., Wu, Y. S., Zheng, J., Wang, Y., Glen, C. R., Collins, D. R., Molina, M. J., and
674 Zhang, R. Y.: Markedly enhanced absorption and direct radiative forcing of black carbon
675 under polluted urban environments, *P. Natl. Acad. Sci. USA*, 113, 4266-4271,
676 doi:10.1073/pnas.1602310113, 2016.

677 Peng, J. F., Hu, M., Wang, Z. B., Huang, X. F., Kumar, P., Wu, Z. J., Guo, S., Yue, D. L.,
678 Shang, D. J., Zheng, Z., and He, L. Y.: Submicron aerosols at thirteen diversified sites in

679 China: size distribution, new particle formation and corresponding contribution to cloud
680 condensation nuclei production, *Atmos. Chem. Phys.*, 14, 10249-10265,
681 doi:10.5194/acp-14-10249-2014, 2014.

682 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic
683 growth and cloud condensation nucleus activity, *Atmos. Chem. Phys.*, 7, 1961-1971,
684 2007.

685 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic
686 growth and cloud condensation nucleus activity—Part 2: Including solubility, *Atmos.*
687 *Chem. Phys.*, 8, 6273– 6279, 2008.

688 Qiu, C., Khalizov, A. F., and Zhang, R. Y.: Soot Aging from OH-Initiated Oxidation of
689 Toluene, *Environ. Sci. Technol.*, 46, 9464-9472, doi:10.1021/Es301883y, 2012.

690 Qiu, C., and Zhang, R.: Multiphase chemistry of atmospheric amines, *Phys. Chem. Chem.*
691 *Phys.* 15, doi:10.1039/C3CP43446J, 5738-5752, 2013.

692 Reed, R. C.: Observations of secondary organic aerosol production and soot aging under
693 atmospheric conditions using a novel new environmental aerosol chamber, PhD
694 dissertation, Texas A&M Univ, College Station, TX., 2010.

695 Ruehl, C. R., Davies, J. F., and Wilson, K. R.: An interfacial mechanism for cloud droplet
696 formation on organic aerosols, *Science*, 351, 1447-1450, doi:10.1126/science.aad4889,
697 2016.

698 Saathoff, H., Naumann, K. H., Schnaiter, M., Schöck, W., Möhler, O., Schurath, U.,
699 Weingartner, E., Gysel, M., and Baltensperger, U.: Coating of soot and (NH₄)₂SO₄
700 particles by ozonolysis products of α -pinene, *J. Aerosol Sci.*, 34, 1297-1321,
701 doi:10.1016/s0021-8502(03)00364-1, 2003.

702 Santoro, R. J., Semerjian, H. G., and Dobbins, R. A.: Soot Particle Measurements in
703 Diffusion Flames, *Combust. Flame*, 51, 203-218, 1983.

704 Swietlicki, E., Hansson, H. C., HÄMeri, K., Svenningsson, B., Massling, A., McFiggans, G.,
705 McMurry, P. H., PetÄJÄ, T., Tunved, P., Gysel, M., Topping, D., Weingartner, E.,
706 Baltensperger, U., Rissler, J., Wiedensohler, A., and Kulmala, M.: Hygroscopic properties
707 of submicrometer atmospheric aerosol particles measured with H-TDMA instruments in
708 various environments—a review, *Tellus B*, 60, 432-469,
709 doi:10.1111/j.1600-0889.2008.00350.x, 2008.

710 Tritscher, T., Jurányi, Z., Martin, M., Chirico, R., Gysel, M., Heringa, M. F., DeCarlo, P. F.,
711 Sierau, B., Prévôt, A. S. H., Weingartner, E., and Baltensperger, U.: Changes of
712 hygroscopicity and morphology during ageing of diesel soot, *Environ. Res. Lett.*, 6,
713 034026, doi:10.1088/1748-9326/6/3/034026, 2011.

714 Wang, G. H., Zhang, R. Y., Gomez, M. E., Yang, L. X., Zamora, M. L., Hu, M., Lin, Y., Peng,
715 J. F., Guo, S., Meng, J. J., Li, J. J., Cheng, C. L., Hu, T. F., Ren, Y. Q., Wang, Y. S., Gao, J.,
716 Cao, J. J., An, Z. S., Zhou, W. J., Li, G. H., Wang, J. Y., Tian, P. F., Marrero-Ortiz, W.,
717 Secret, J., Du, Z. F., Zheng, J., Shang, D. J., Zeng, L. M., Shao, M., Wang, W. G., Huang,
718 Y., Wang, Y., Zhu, Y. J., Li, Y. X., Hu, J. X., Pan, B., Cai, L., Cheng, Y. T., Ji, Y. M.,
719 Zhang, F., Rosenfeld, D., Liss, P. S., Duce, R. A., Kolb, C. E., and Molina, M. J.:
720 Persistent sulfate formation from London Fog to Chinese haze, *P. Natl. Acad. Sci. USA*,
721 113, 13630-13635, doi:10.1073/pnas.1616540113, 2016.

722 Wang, L., Khalizov, A. F., Zheng, J., Xu, W., Ma, Y., Lal, V., and Zhang, R. Y.: Atmospheric

723 nanoparticles formed from heterogeneous reactions of organics, *Nat. Geosci.*, 3, 238-242,
724 doi:10.1038/NGEO778, 2010.

725 Wang, Y., Wan, Q., Meng, W., Liao, F., Tan, H., and Zhang, R.: Long-term impacts of
726 aerosols on precipitation and lightning over the Pearl River Delta megacity area in China,
727 *Atmos. Chem. Phys.* 11, 12421–12436, 2011.

728 Wang, Y., Khalizov, A., Levy, M., and Zhang, R.: Light absorbing aerosols and their
729 atmospheric impacts, *Atmos. Environ.* 81, 713-715, doi:10.1016/j.atmosenv.2013.09.034,
730 2013.

731 Weingartner, E., Burtscher, H., and Baltensperger, U.: Hygroscopic properties of carbon and
732 diesel soot particles, *Atmos. Environ.*, 31, 2311-2327,
733 doi:10.1016/S1352-2310(97)00023-X, 1997.

734 Wittbom, C., Eriksson, A. C., Rissler, J., Carlsson, J. E., Roldin, P., Nordin, E. Z., Nilsson, P.
735 T., Swietlicki, E., Pagels, J. H., and Svenningsson, B.: Cloud droplet activity changes of
736 soot aerosol upon smog chamber ageing, *Atmos. Chem. Phys.*, 14, 9831-9854,
737 doi:10.5194/acp-14-9831-2014, 2014.

738 Wu, Z. J., Zheng, J., Shang, D. J., Du, Z. F., Wu, Y. S., Zeng, L. M., Wiedensohler, A., and
739 Hu, M.: Particle hygroscopicity and its link to chemical composition in the urban
740 atmosphere of Beijing, China, during summertime, *Atmos. Chem. Phys.*, 16, 1123-1138,
741 doi:10.5194/acp-16-1123-2016, 2016.

742 Xue, H., Khalizov, A. F., Wang, L., Zheng, J., and Zhang, R.: Effects of dicarboxylic acid
743 coating on the optical properties of soot, *Phys. Chem. Chem. Phys.*, 11, 7865-7875,
744 doi:10.1039/b700001a, 2009a.

745 Xue, H., Khalizov, A. F., and Zhang, R.: Effects of coating of dicarboxylic acids on the
746 mass-mobility relationship of soot particles, *Environ. Sci. Technol.* 43, 2787–2792, doi:
747 10.1021/es803287v, 2009b.

748 Yuan, T., Li, Z., Zhang, R., and Fan, J.: Increase of cloud droplet size with aerosol optical
749 depth: An observation and modeling study, *J. Geophys. Res.*, 113, D04201,
750 doi:10.1029/2007JD008632, 2008.

751 Zhang, D., and Zhang R.: Laboratory investigation of heterogeneous interaction of sulfuric
752 acid with soot, *Environ. Sci. Technol.* 39, 5722-5727, doi:10.1021/es050372d, 2005.

753 Zhang, R., Khalizov, A. F., Pagels, J., Zhang, D., Xue, H., and McMurry, P. H.: Variability in
754 morphology, hygroscopicity, and optical properties of soot aerosols during atmospheric
755 processing, *P. Natl. Acad. Sci. USA*, 105, 10291-10296, doi:10.1073/pnas.0804860105,
756 2008.

757 Zhang, R. Y., Suh, I., Zhao, J., Zhang, D., Fortner, E. C., Tie, X. X., Molina, L. T., and
758 Molina, M. J.: Atmospheric new particle formation enhanced by organic acids, *Science*,
759 304, 1487-1490, doi:10.1126/science.1095139, 2004.

760 Zhang, R., Khalizov, A.F., Wang, L., Hu, M., Xu, W.: Nucleation and growth of nanoparticles
761 in the atmosphere, *Chem. Rev.*, 112, 1957-2011, doi:10.1021/cr2001756, 2012.

762 Zhang, R. Y., Wang, G. H., Guo, S., Zarnora, M. L., Ying, Q., Lin, Y., Wang, W. G., Hu, M.,
763 and Wang, Y.: Formation of Urban Fine Particulate Matter, *Chem. Rev.*, 115, 3803-3855,
764 doi:10.1021/acs.chemrev.5b00067, 2015.

765 Zhao, J., Levitt, N. P., Zhang, R., and Chen, J.: Heterogeneous reactions of methylglyoxal in
766 acidic media: Implications for secondary organic aerosol formation, *Environ. Sci.*

767 Technol., 40, 7682-7687, doi:10.1021/es060610k, 2006.
768 Zhao, J., Khalizov, A.F., Zhang, R., and McGraw, R.: Hydrogen bonding interaction of
769 molecular complexes and clusters of aerosol nucleation precursors, *J. Phys. Chem.*, 113,
770 680–689, doi:10.1021/jp806693r, 2009.
771 Zhao, P. S., Dong, F., He, D., Zhao, X. J., Zhang, X. L., Zhang, W. Z., Yao, Q., and Liu, H. Y.:
772 Characteristics of concentrations and chemical compositions for PM_{2.5} in the region of
773 Beijing, Tianjin, and Hebei, China, *Atmos. Chem. Phys.*, 13, 4631-4644,
774 doi:10.5194/acp-13-4631-2013, 2013.
775 Zheng, J., Zhang, R., Fortner, E. C., Volkamer, R. M., Molina, L., Aiken, A. C., Jimenez, J. L.,
776 Gaeggeler, K., Dommen, J., Dusanter, S., Stevens, P.S., and Tie, X.: Measurements of
777 HNO₃ and N₂O₅ using ion drift - chemical ionization mass spectrometry during the
778 MCMA - 2006 campaign, *Atmos. Chem. Phys.*, 8, 6823–6838, 2008.
779 Zheng, J., Khalizov, A., Wang, L., and Zhang, R. Y.: Atmospheric Pressure-Ion Drift
780 Chemical Ionization Mass Spectrometry for Detection of Trace Gas Species, *Anal. Chem.*,
781 82, 7302-7308, doi:10.1021/ac101253n, 2010.

Table 1. Summary of ambient conditions for BC aging experiments conducted in Beijing. The PM_{2.5}, gas concentrations and meteorological condition were averaged from the entire experimental time. D_m and T represent the mobility diameter and temperature, respectively.

No.	Date	Time		Initial D _m (nm)	PM _{2.5} (μg m ⁻³)	J(O ¹ D) (10 ⁻⁶)	Gas Concentration (ppb)					Meteorological Conditions	
		Start	End				Toluene	Xylene	SO ₂	NO _x	O ₃	T (°C)	RH (%)
#1	Aug. 18 th	12:41	14:47	95	43	19	0.49	0.13	1.7	8.8	56	35	27
#2	Aug. 22 nd	13:32	16:51	96	69	3.7	3.41	0.78	2	36.1	26	26	69
#3	Sep. 7 th	12:40	14:52	97	12	17.5	0.71	0.17	2.7	10.2	75	30	35
#4	Sep. 9 th	13:13	17:06	97	40	6.3	0.77	0.21	4	11	68	26	50
#5	Sep. 1 st	13:19	16:13	147	27	11.6	0.76	0.19	4.6	19.9	83	31	33
#6	Sep. 11 st	13:50	17:25	147	57	6.1	1.57	0.39	6.7	17.1	92	29	42
#7	Sep. 21 st	15:31	17:41	146	30	2.1	0.75	0.29	2	10.6	90	28	37
#8	Aug. 24 th	11:37	16:06	216	8.8	21.1	0.98	0.3	1.7	15.6	57	36	25
#9	Sep. 5 th	14:06	16:44	220	12	8.3	0.45	0.2	2.2	14.6	54	29	35
#10	Oct. 17 th	12:54	17:13	224	57	3.2	-	-	13.8	41	34	18	30

Table 2. Summary of particle properties for BC aging experiments conducted in Beijing. D_m , ρ_{eff} and D_{me} represent the mobility diameter, effective density and mass equivalent diameter, respectively.

No.	Date	D_m		ρ_{eff}		D_{me}		ΔD_{me} (nm)	Growth Rate ^a (nm h ⁻¹)
		Initial (nm)	Final (nm)	Initial (g cm ⁻³)	Final (g cm ⁻³)	Initial (nm)	Final (nm)		
#1	Aug. 18 th	95	157	0.50	1.35	62	162	100	47
#2	Aug. 22 nd	96	129	0.46	1.31	61	126	65	20
#3	Sep. 7 th	97	147	0.45	1.25	62	142	80	36
#4	Sep. 9 th	97	136	0.43	1.30	61	133	62	19
#5	Sep. 1 st	147	170	0.34	1.36	85	168	83	29
#6	Sep. 11 st	147	162	0.34	1.34	84	159	75	20
#7	Sep. 21 st	146	132	0.34	1.05	84	116	32	15
#8	Aug. 24 th	216	272	0.32	1.37	123	275	152	34
#9	Sep. 5 th	220	202	0.25	1.33	114	197	83	31
#10	Oct. 17 th	224	224	0.24	0.52	117	157	40	11
Average								77 ± 33	26 ± 11

^a The growth rate is calculated using the data between 12:00 and 17:00 for each experiment.

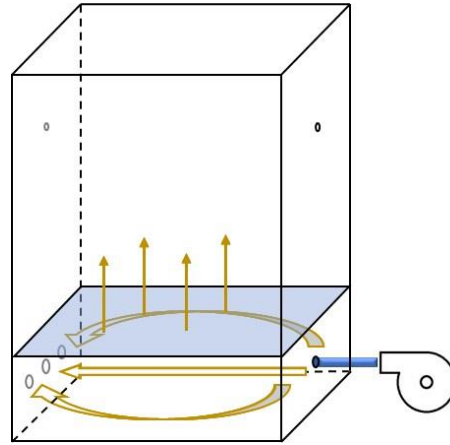


Figure 1. Photo (left) and schematic (right) of the quasi-atmospheric aerosol evolution study (QUALITY) chamber

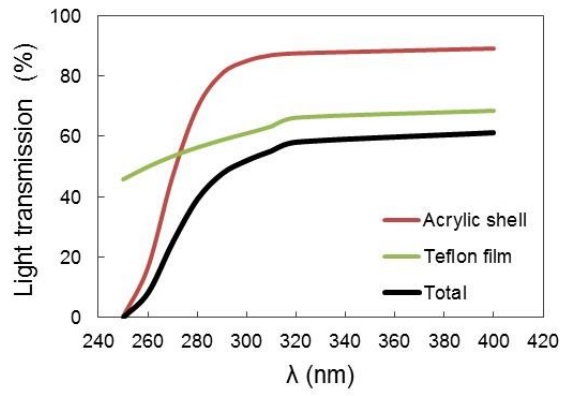


Figure 2. Light transmission spectra of the PFA Teflon film (yellow line), the acrylite shell (red line), and their total transmission in the UV range (black line).

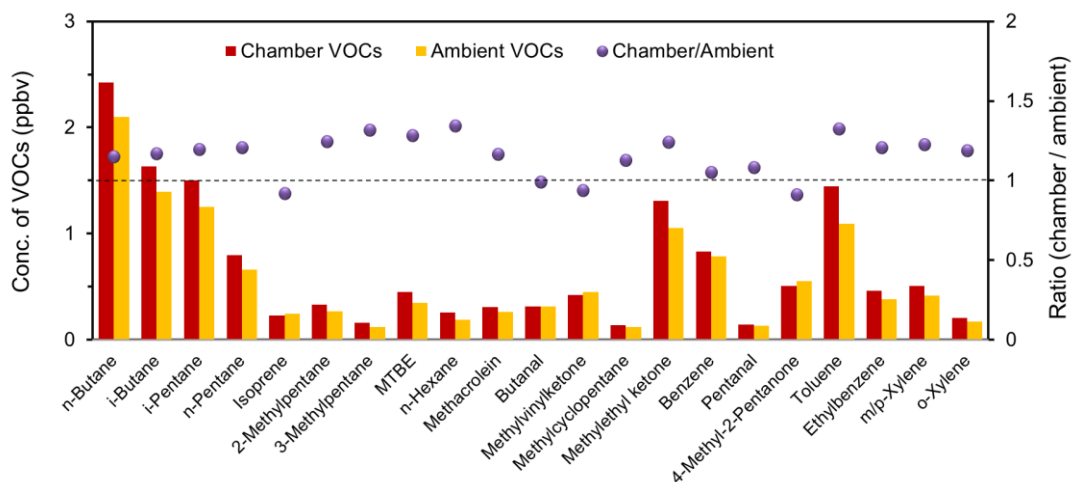


Figure 3. Comparison of VOCs concentrations in the QUALITY chamber (red bar) and in the ambient air (yellow bar) before the start of BC aging experiment. Purple cycles represent the ratio of each species in the reaction chamber to in the ambient air. Only the VOCs species that contains 4 or more carbons with the concentration higher than 0.1 ppb are shown in the figure.

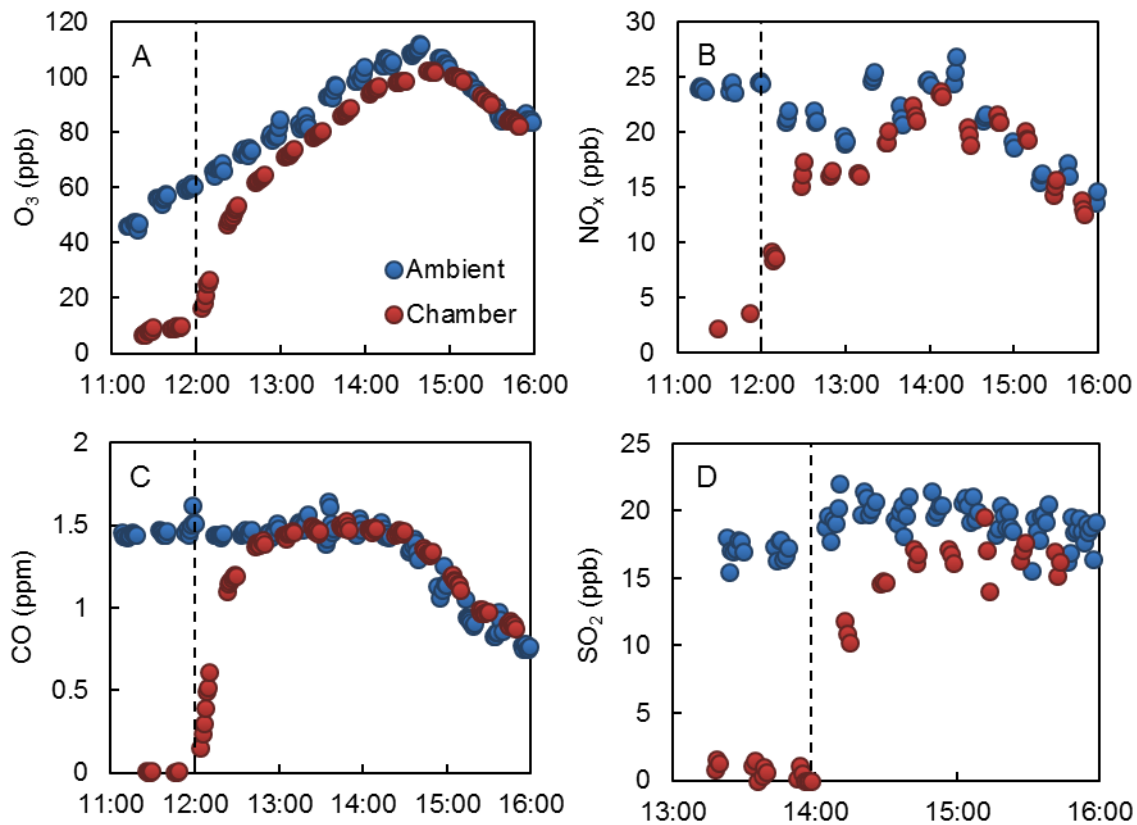


Figure 4. Concentrations of O₃ (A), NO_x (B), CO (C), and SO₂ (D) measured inside the QUALITY chamber (Red circles) and in the ambient air (blue circles). The vertical dashed lines denote the time when the ambient air started to be pulled through the bottom flow chamber and the ambient gases began to exchange into the upper reaction chamber.

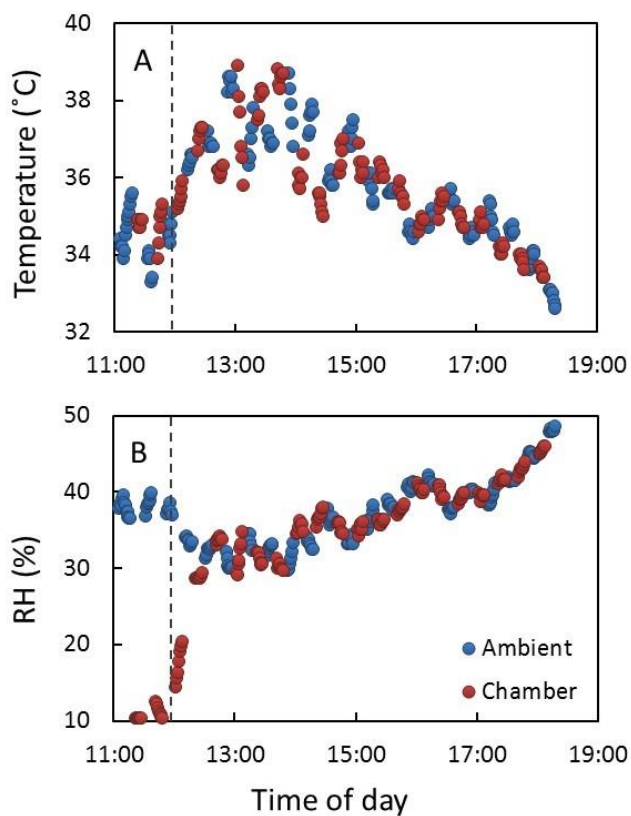


Figure 5. Temperature (A) and RH (B) measured inside the QUALITY chamber (Red circles) and in the ambient air (blue circles). The temperature and RH were measured by probes placed in the sampling tube adjacent to the chamber. The vertical dashed lines denote the time when the ambient air started to be pulled through the bottom flow chamber and the ambient gases began to exchange into the upper reaction chamber.

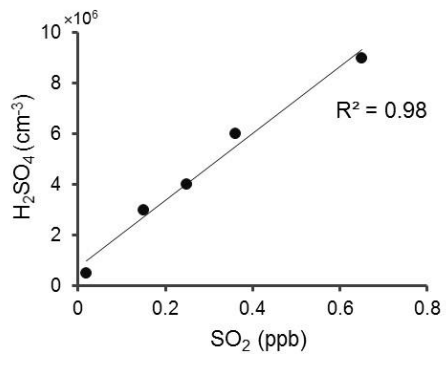


Figure 6. Sulfuric acid concentration as function of SO₂ concentration inside the chamber

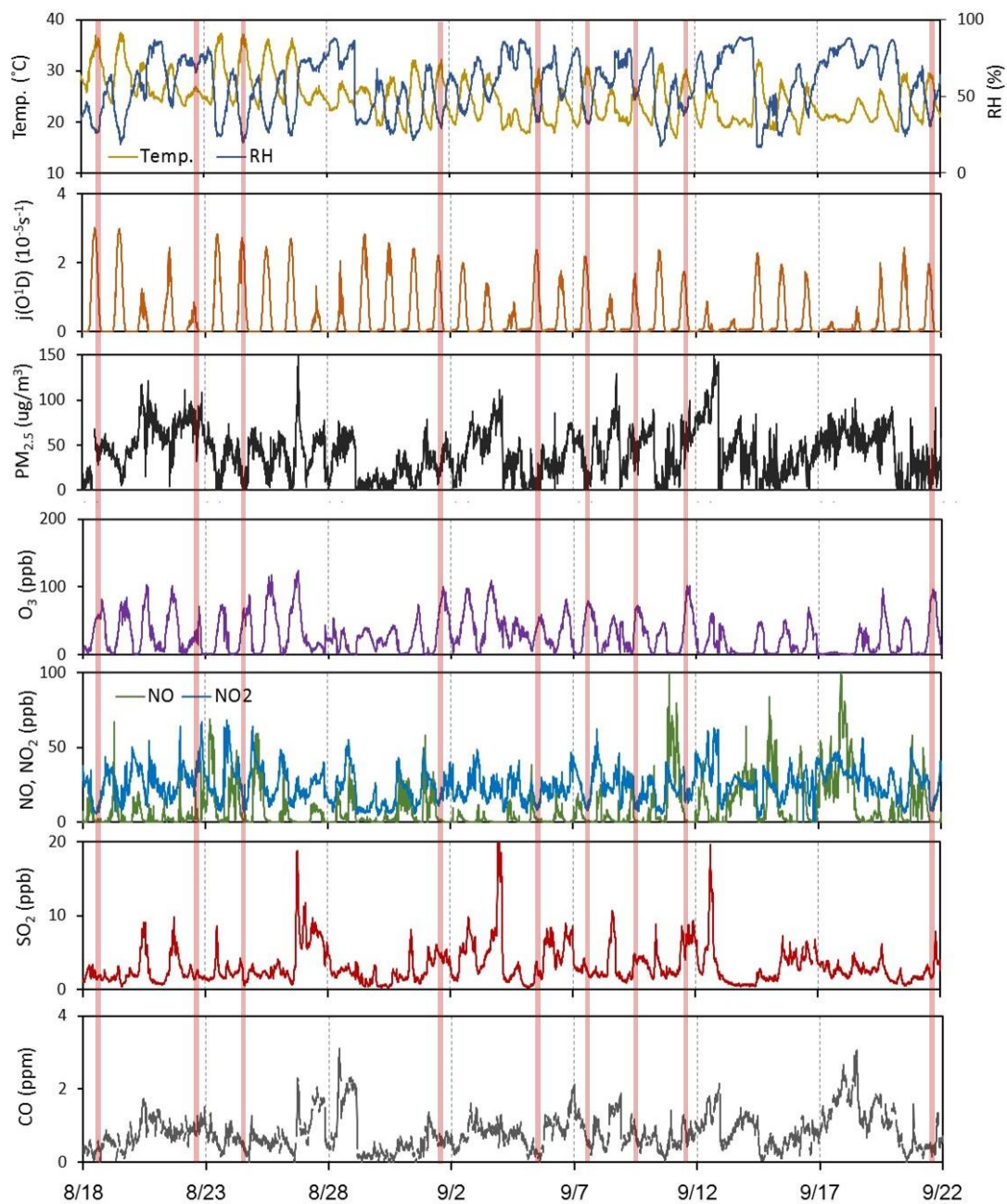


Figure 7. Time series of ambient pollutant concentrations and meteorological parameters during the experimental period in Beijing. Red bars indicate periods when BC aging experiments were conducted.

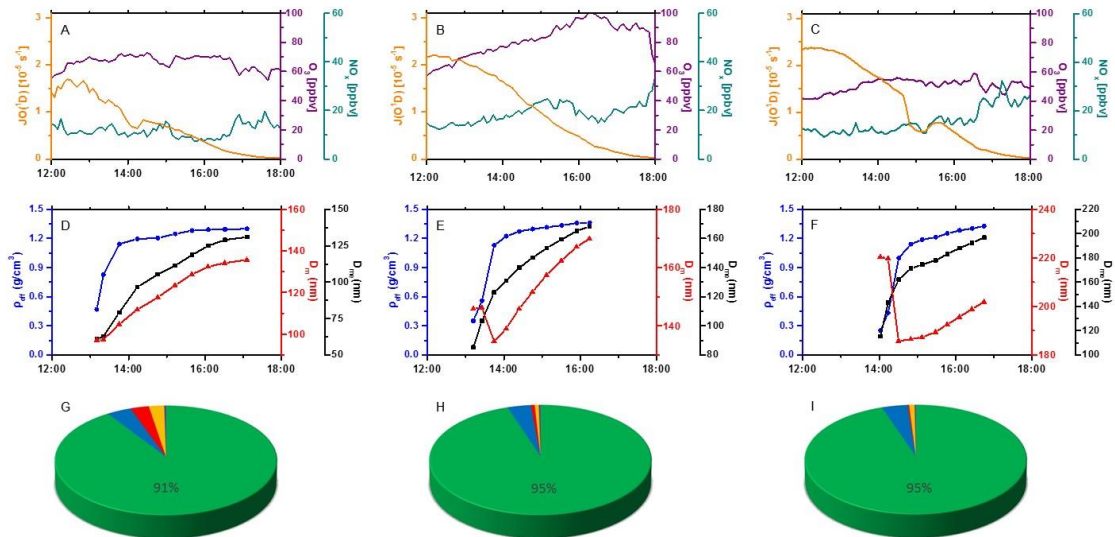


Figure 8. Ambient condition (A, B, C), changes of diameter and density of BC particle (D, E, F), and the chemical composition of coating materials (G, H, I) during three typical aging experiments. A and D correspond to experiment #4; B and E correspond to experiment #5, and C and E correspond to experiment #9. D_m is the peak mobility diameter of BC particles, ρ_{eff} is the best-fit effective density of BC particles with the mobility diameter D_m , D_{me} is the mass equivalent diameter of BC particles, and $J(O^1D)$ represents the measured photolysis rate constant for $O(^1D)$. The colors of green, blue, red, yellow and purple in the pie charts represent organics, nitrate, sulfate, ammonium, and chlorine, respectively. The numbers in figure G, H and I are the mass fraction of organics.

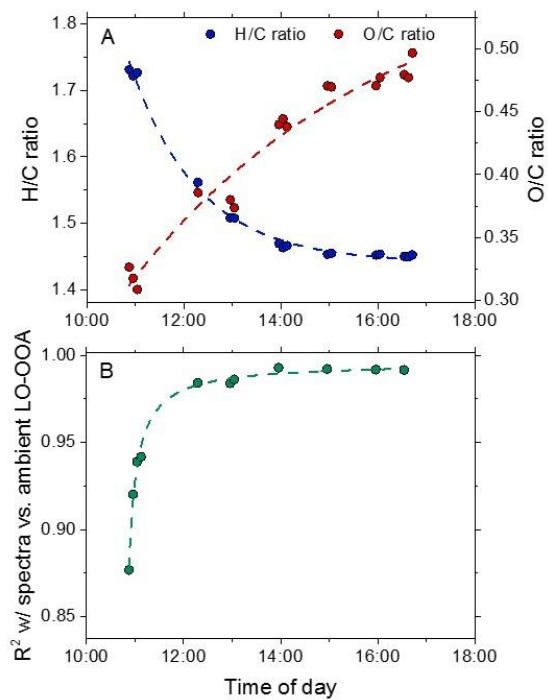


Figure 9. The evolution of organic aerosols inside the chamber during an aging experiment (#8) in Beijing. (A) the H/C and O/C ratios of organics on aged BC particles; (b) the correlation coefficients (R^2) between the evolving total OA spectra in chamber experiment and the LO-OOA spectra derived from the Beijing field data set.

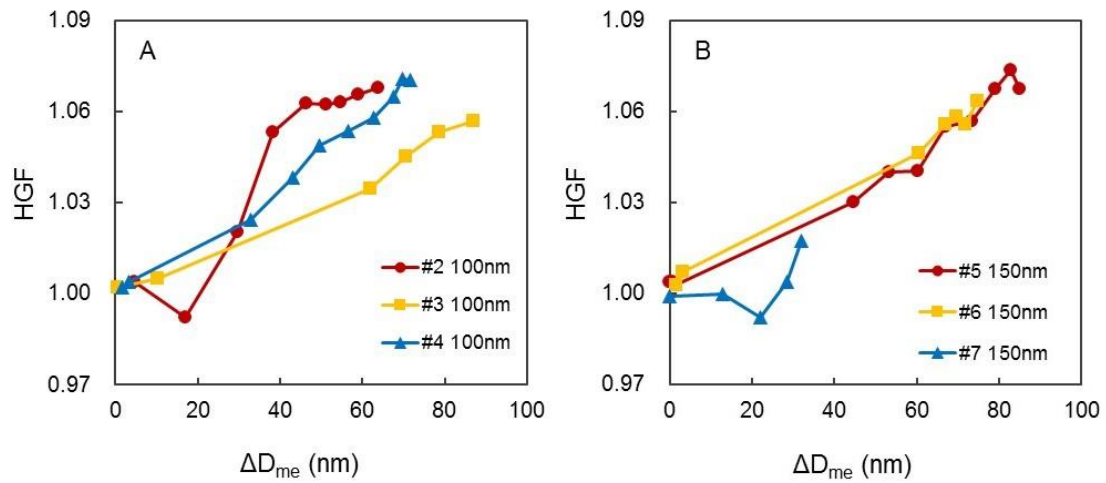


Figure 10. Evolution of hygroscopic growth factors (HGF) of BC particles during aging as a function of ΔD_{me} . (A) three experiments with 100 nm BC particle; (B) three experiments with 150 nm BC particle. Different colors in each figure represent different experiments. Hygroscopicity measurement is not available for experiment #1.

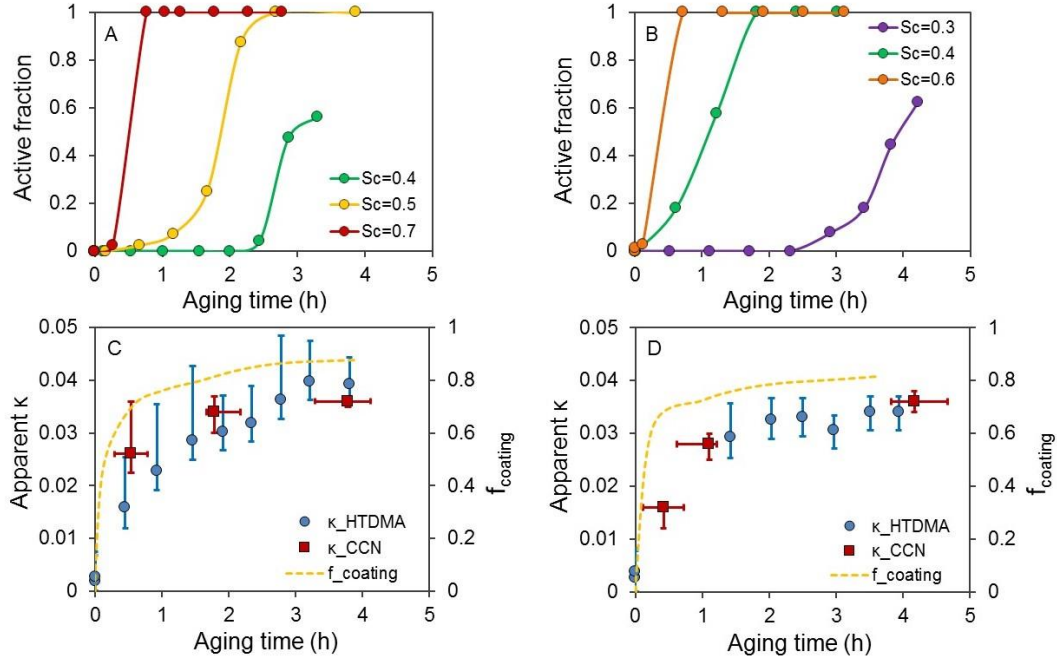


Figure 11. The active fraction of BC particles under diversified supersaturation (A, B) and the closure of apparent κ for BC particles with initial diameter (C, D) during aging in two typical experiments. A, C are the results from experiment #4 with 100nm BC particles and B, D represents experiment #6 with 150nm BC particles. Red square and blue circle in C and D represent the apparent κ calculated using CCN counter data and HTDMA data, respectively. Yellow slash line represents the fraction of coating materials on BC particles. The error bars of κ_{CCN} and κ_{HTDMA} represent the uncertainty in the calculation.