

Dear Editor,

Please find enclosed the revised version of our manuscript entitled "IASI-derived NH_3 enhancement ratios relative to CO for the tropical biomass burning regions". We have addressed all the comments of the referees.

We would also like to draw your attention to the following. After the initial submission of the manuscript, a major update in the IASI- NH_3 retrieval product was introduced. This update is of importance especially since we identified in the previous version of the NH_3 dataset sharp discontinuities in the NH_3 time series related to inconsistencies between the different versions of the operational Eumetsat IASI L2 algorithms for temperature and clouds (which are used as input in the neural network). To tackle this issue, we have developed a version of the IASI- NH_3 Neural Network which relies on the ERA-Interim ECMWF meteorological input data (along with built-in surface temperature) rather than on the Eumetsat IASI L2 data. This provides a reanalyzed dataset which is coherent in time over the whole period covered by IASI (2008-today) and can therefore be used to investigate interannual variability. The reanalyzed dataset has been described in a paper recently published in AMTD¹.

After careful analysis, we have judged that this new dataset should ideally be used also for the present manuscript. We have updated all the figures of the manuscript according to this new dataset. As you will see, the main conclusions of our paper remain unchanged, except for the comparison with the $\text{ER}_{\text{NH}_3/\text{CO}}$ from the literature. The IASI derived $\text{ER}_{\text{NH}_3/\text{CO}}$ are higher than in the manuscript initially submitted and this leads to a much better agreement with the values given in the literature. With these changes and the comments of the referees addressed, we are confident that the paper has been greatly improved.

The responses to each of the referee's comments are provided here below. A 'tracked changes' version of the manuscript is appended and allows to better identify where and how the manuscript was modified. We hope that with these changes and the replies made to the referees, you will find the manuscript suitable for publication in ACP.

Yours sincerely,

Simon Whitburn.

¹ Van Damme, M., Whitburn, S., Clarisse, L., Clerbaux, C., Hurtmans, D., and Coheur, P.-F.: Version 2 of the IASI NH_3 neural network retrieval algorithm; near-real time and reanalysed datasets, Atmos. Meas. Tech. Discuss., <https://doi.org/10.5194/amt-2017-239>, in review, 2017.

Response to reviewers:

Reviewer #1:

We would like to thank the reviewer for his/her positive feedback on the paper and useful comments. All have been addressed. The point-by-point responses to each of the comments are provided below. A 'tracked changes' version of the manuscript is also appended.

We would also like to draw your attention to the following. After the initial submission of the manuscript, a major update in the IASI-NH₃ retrieval product was introduced. This update is of importance especially since we identified in the previous version of the NH₃ dataset sharp discontinuities in the NH₃ time series related to inconsistencies between the different versions of the operational Eumetsat IASI L2 algorithms for temperature and clouds (which are used as input in the neural network). To tackle this issue, we have developed a version of the IASI-NH₃ Neural Network which relies on the ERA-Interim ECMWF meteorological input data (along with built-in surface temperature) rather than on the Eumetsat IASI L2 data. This provides a reanalyzed dataset which is coherent in time over the whole period covered by IASI (2008-today) and can therefore be used to investigate interannual variability. The reanalyzed dataset has been described in a paper recently published in AMTD².

After careful analysis, we have judged that this new dataset should ideally be used also for the present manuscript. We have updated all the figures of the manuscript according to this new dataset. As you will see, the main conclusions of our paper remain unchanged, except for the comparison with the ER_{NH₃/CO} from the literature. The IASI derived ER_{NH₃/CO} are higher than in the manuscript initially submitted and this leads to a much better agreement with the values given in the literature. With these changes and the comments addressed, we are confident that the paper has been greatly improved.

3_30: The expression 'fair' is only qualitative. Please describe the agreement in more quantitative terms.

We have changed the sentence in the manuscript to:

Two studies, based on a previous NH₃ retrieval algorithm also using the HRI but relying on two-dimensional look-up tables for the conversion into a NH₃ total column (molec.cm⁻²) (Van Damme et al., 2014), have shown a fair agreement between IASI-NH₃ observations and other measurements (generally within the uncertainties of the IASI NH₃ retrieved columns), with differences of about 60-80% reported in Van Damme et al. (2015) and of 30% on average in Dammers et al. (2016).

² Van Damme, M., Whitburn, S., Clarisse, L., Clerbaux, C., Hurtmans, D., and Coheur, P.-F.: Version 2 of the IASI NH₃ neural network retrieval algorithm; near-real time and reanalysed datasets, Atmos. Meas. Tech. Discuss., <https://doi.org/10.5194/amt-2017-239>, in review, 2017.

4_1: "We also have assumed a similar sensitivity for IASI to NH₃ and CO in the lower layers of the atmosphere, which is not expected to introduce a significant bias in the studied regions due to a generally positive thermal contrast prevailing during daytime".

- Can you demonstrate this assumption e.g. by presenting typical averaging kernels showing such a similar sensitivity of the NH₃ and CO retrievals at the lower altitudes?

The retrieval algorithm for NH₃ does not provide averaging kernels. It is therefore difficult to compare quantitatively the sensitivity of NH₃ and CO in the lower layers. However, the sensitivity of thermal nadir measurements near the surface is intimately related to the thermal contrast between the surface and the first layers of the atmosphere (Clerbaux et al., 2009). For NH₃, Clarisse et al. (2010) have shown that, when the detection is possible (i.e. in case of good thermal contrast), the peak sensitivity for NH₃ is in the boundary layer. For CO, Georges et al. (2009) and Bauduin et al. (2016) have also demonstrated the good sensitivity of IASI in the lower layers in case of high thermal contrast. Since we consider here only daytime measurements, in tropical regions and with a retrieval uncertainty on the NH₃ column lower than 100% (i.e. a good sensitivity), the majority of the observations retained correspond to this situation of good thermal contrast.

George, M., Clerbaux, C., Hurtmans, D., Turquety, S., Coheur, P.-F., Pommier, M., Hadji-Lazaro, J., Edwards, D. P., Worden, H., Luo, M., Rinsland, C., and McMillan, W. (2009). Carbon monoxide distributions from the IASI/METOP mission: evaluation with other space-borne remote sensors. *Atmos. Chem. Phys.*, 9(21):8317–8330. doi: 10.5194/acp-9-8317-2009.

Bauduin, S. and Clarisse, L. and Theunissen, M. and George, M. and Hurtmans, D. and Clerbaux, C. and Coheur, P. F. (2016). IASI's sensitivity to near-surface carbon monoxide (CO): Theoretical analyses and retrievals on test cases. *Journal of Quantitative Spectroscopy & Radiative Transfer*, 189:428-440. doi: /10.1016/j.jqsrt.2016.12.022.

To clarify this, we have adapted the following sentence in the manuscript to:

We also have assumed a similar sensitivity for IASI to NH₃ and CO in the lower layers of the atmosphere. This is not expected to introduce a significant bias since it has been shown for both CO and NH₃ that the peak sensitivity was in the lower layers of the atmosphere in case of positive thermal contrast which is generally prevailing in the studied regions during daytime (George et al., 2009; Clarisse et al., 2010; Van Damme et al., 2014; Bauduin et al., 2016).

- Due to the different values of thermal contrast between surface and the lower atmosphere for different observations, an error could be introduced in the dataset which might partly be responsible for the overall variability of the ERs. Have you tried to correlate the temperature contrast with your dataset and can you exclude such an influence?

It is true that the thermal contrast greatly influence the sensitivity of IASI to near-surface measurements. However, as mentioned here above, for the observation considered, we are generally in a situation of high positive thermal contrast where NH₃ (and CO) can generally be retrieved with a good accuracy. Moreover, the sensitivity of the retrieved column to the atmospheric parameters (including the thermal contrast) is included in the uncertainty.

4_5: "Whitburn et al. (2016b) have calculated that the use of an alternative profile could affect the retrieved column up to 50%" It would be very helpful if one could try to derive a quantitative error assessment for the NH₃/CO ratios combining estimated errors of NH₃ and CO, taking into account such effect as described. At least it would be very helpful to put together a table where specific uncertainties and their source are listed for NH₃ and CO separately.

Since the $ER_{NH_3/CO}$ are calculated from the slope of the linear regression of NH₃ versus CO total columns, they do not rely on the individual uncertainties of each NH₃ and CO observation. It is therefore very difficult to estimate a quantitative error assessment for the NH₃/CO ratios.

As explained in the next comment, the total uncertainty on the NH₃ retrieved columns depend on the atmospheric conditions and on the amount of NH₃. Examples of the relative importance of the different atmospheric parameters on the total uncertainty are given in Whitburn et al. (2016b). The observations considered in this study correspond generally to a situation of high thermal contrast and a relatively high NH₃ column. In the case of high thermal contrast and high NH₃ column, Whitburn et al. (2016b) have shown that the contribution to the total uncertainty was well distributed on the different parameters while, for high thermal contrast but lower columns, the major contribution to the total uncertainty was on the HRI (mainly instrumental noise). Note that Whitburn et al. (2016b) have shown that, despite a possible high uncertainty, the average bias on the columns is close to zero.

For CO, the uncertainty is generally lower (of about 5-10% and up to 20% for the observations considered here). The uncertainty on the calculated $ER_{NH_3/CO}$ is therefore mainly driven by the uncertainty on the NH₃ column (which is lower than 100% due to the pre-filtering applied).

5_26: "With a relative error lower than 100% for NH₃": Which kind of error is this? Only due to spectral noise?

The error associated with the NH₃ retrieved columns includes a full uncertainty analysis and depend on the atmospheric conditions and the amount of NH₃. The uncertainty analysis is performed by perturbing the input parameters (Temperature profile, NH₃ a priori vertical profile, HRI, etc.) of the neural network and evaluates for each retrieved column how the uncertainty of the input parameters propagates to the final result.

We have added the following sentence in the manuscript to make it more clear for the reader (3_21):

The retrieval also includes a full uncertainty analysis, performed by perturbing the input parameters (temperature profile, HRI, NH_3 a priori profile, etc.) of the neural network.

8_5: "The agreement would become even better if we consider in addition a possible bias due to the use of a non-representative NH_3 vertical profile." What is the reason that the agreement would improve and not becoming worse when using another a-priori NH_3 profile?

Thank you for pointing this out. It is indeed very hard to evaluate whether the 'real' NH_3 profiles would increase or decrease the retrieved columns because of the large variety of NH_3 profiles associated with transported fire plumes. As mentioned above, Clarisse et al. (2010) have shown that in case of favorable thermal contrast, the peak sensitivity of IASI for NH_3 is in the boundary layer (below 1 km). For the fire plumes, Val Martin et al. (2010) have shown that, for tropical forests and grasslands/croplands fires, the plume height was generally below the boundary layer height while only a small fraction peaked above 1.5 km. Due to a lack of available information on plume height for each individual fires, it is hard to determine the impact of the true profile on the retrieved columns (and therefore on the ratio NH_3/CO).

M. Val Martin, J. A. Logan, R. A. Kahn, F.-Y. Leung, D. L. Nelson, and D. J. Diner. Smoke injection heights from fires in North America: analysis of 5 years of satellite observations. *Atmospheric Chemistry and Physics*, 10(4):1491-1510, 2010. doi: 10.5194/acp-10-1491-2010.

We have adapted the following sentence to clarify this:

Finally, as we mentioned in section 2.1, the retrieval of NH_3 could be biased by the use of a constant NH_3 vertical profile not representative of the variety of profiles observed above biomass burning plumes.

10_8: "These differences may be explained by various factors including...": In the list, possible errors regarding the NH_3 (and CO) IASI retrievals is missing though it has been mentioned before.

We have modified the following sentence to take this into account:

This may be explained by various factors including 1) the parametrization (pre- and post-filtering of the data) considered for the calculation of the $\text{ER}_{\text{NH}_3/\text{CO}}$, 2) a bias towards the flaming phase due to the selection of IASI observations close to MODIS active fires (less sensitive to the smoldering phase) and 3) a possible accumulation of CO in the region during the fire season, introducing a low-bias in the IASI-derived $\text{ER}_{\text{NH}_3/\text{CO}}$. Another possible explanation might lie in the use of a unique vertical profile shape in the retrieval scheme of NH_3 while biomass burning plumes exhibit a large variety of plume injection heights.

23_Table1: It would be very helpful if you could list, e.g. after each number, the result from the actual analysis for a similar biome and geographic region.

This is indeed a good suggestion. We have added to table 1 two rows including the results from the actual analysis (one on a regional level and one for a biome level).

Technical comments have been taken into account.

Reviewer #2:

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The relationship to NO₂. Are the NH₃ and NO₂ columns related?

NO₂ and NH₃ emissions are anti-correlated during the different stages of a fire. During the flaming phase, which takes place first, emissions of oxidized species (including NO₂) dominate. When all the easily accessible fuel has been consumed, smoldering combustion starts with increased emissions of reduced or incompletely oxidized species (NH₃, CO, etc.) due to the lack of available atmospheric oxygen and/or the lower temperatures.

Flaming vs. smoldering fires might have quite different CO but similar NH₃. Is anything known? The paper gives the impression that CO is more constant than NH₃. I don't know if

³ Van Damme, M., Whitburn, S., Clarisse, L., Clerbaux, C., Hurtmans, D., and Coheur, P.-F.: Version 2 of the IASI NH₃ neural network retrieval algorithm; near-real time and reanalysed datasets, Atmos. Meas. Tech. Discuss., <https://doi.org/10.5194/amt-2017-239>, in review, 2017.

that is true, but I think the paper should comment on whether variability is driven more by CO changes or more by the N-content of fuels. For example, wood has zero N, a recent paper Coggon et al. Geophys. Res. Lett. 2016 pointed out that residential wood burning has near zero emissions of N-compounds.

Flaming and smoldering phases show very different proportions of NH_3 emissions relative to the total Nr emissions. As mentioned above, emissions of oxidized Nr species (mainly NO_2) dominate during the flaming phase while, during the smoldering phase, NH_3 becomes the dominant Nr species emitted. CO also dominates during the smoldering phase.

This is indeed true (and well known) that flaming and smoldering phases will have very different $\text{ER}_{\text{NH}_3/\text{CO}}$. This is observed in particular in this study when looking at the differences between the $\text{ER}_{\text{NH}_3/\text{CO}}$ for tropical forest and the savanna fires. Forest fires, which are characterized by a larger fraction of smoldering combustion than savannas (due to the higher biomass density) show higher $\text{ER}_{\text{NH}_3/\text{CO}}$ than savannas. When studying enhancement ratios from satellite observations, we calculate an average $\text{ER}_{\text{NH}_3/\text{CO}}$ (of generally an entire fire) over a certain time period and area which therefore include both smoldering and fire phases. Resulting average $\text{ER}_{\text{NH}_3/\text{CO}}$ is therefore the result of many different factors acting together. It is however not possible to separate the contribution of each ones.

IASI-derived NH_3 enhancement ratios relative to CO for the tropical biomass burning regions

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Abstract. Vegetation fires are a major source of ammonia (NH_3) in the atmosphere. Their emissions are mainly estimated from "bottom-up" approaches which rely on uncertain emission factors. In this study, we derive new biome-specific NH_3 enhancement ratios relative to carbon monoxide (CO), $\text{ER}_{\text{NH}_3/\text{CO}}$ – directly related to the emission factors, from the measurements of the IASI sounder on board the Metop-A satellite. This is achieved for large tropical regions and for a 8-year period (2008–2015). We find substantial differences in the $\text{ER}_{\text{NH}_3/\text{CO}}$ between the studied biomes with calculated values ranging from 4.47×10^{-3} to 1723×10^{-3} . For Evergreen Broadleaf Forest these are typically 75–100% higher than for Woody Savanna and Savanna biomes. This variability is attributed to differences in fuel types and size and is in line with previous studies. The analysis of the spatial and temporal distribution of the $\text{ER}_{\text{NH}_3/\text{CO}}$ also reveals a (sometimes large) within-biome variability. On a regional level, Woody Savanna shows for example a mean $\text{ER}_{\text{NH}_3/\text{CO}}$ for the region of Africa South of the Equator which is 50–100% lower than in the other five studied regions, probably reflecting regional differences in fuel type and burning conditions. The same variability is also observed on a yearly basis with a peak in the $\text{ER}_{\text{NH}_3/\text{CO}}$ observed for the year 2010 for all biomes. These results highlight the need for the development of dynamic emission factors that better take into account local variations in fuel type and fire conditions. We also compare the IASI-derived $\text{ER}_{\text{NH}_3/\text{CO}}$ with values reported in the literature, usually calculated from ground-based or airborne measurements. We find a general underestimation over the referenced $\text{ER}_{\text{NH}_3/\text{CO}}$ of about 40% for Woody Savanna and Savanna and up to a factor 1.5–4 for Evergreen Broadleaf Forest and Cropland. Beyond a possible overestimation of except for Cropland for which the $\text{ER}_{\text{NH}_3/\text{CO}}$ in the literature, the observed differences could also be related to various factors including instrumental limits, bias in the retrieval of the NH_3 columns, parameterization in the calculation of the $\text{ER}_{\text{NH}_3/\text{CO}}$ or accumulation of CO in the studied regions during the fire period. show an underestimation of about 2–2.5 times.

1 Introduction

Vegetation fires contribute significantly to the global budget of many trace gases and aerosols in the atmosphere (Langmann et al., 2009). Carbon dioxide (CO_2) emissions from biomass burning are for example estimated to about 2–4 Pg C year^{−1} against 7.2 Pg C year^{−1} from fossil fuel combustion (Bowman et al., 2009). For carbon monoxide (CO), the contribution

to the total budget could even reach more than 50% (Crutzen and Andreae, 1990; van der Werf et al., 2004, 2010). Besides carbon, vegetation fires also emit large amounts of reactive nitrogen species, among which ammonia (NH_3). With a contribution estimated to be about 13% (Galloway et al., 2004) of the total emissions, biomass burning is believed to be the second most important source of NH_3 after agriculture. From previous studies, it has been shown that biomass burning could affect significantly NH_3 concentrations in the atmosphere, especially in the tropics but also at higher latitudes (e.g., Bouwman et al., 1997; Coheur et al., 2009; Adon et al., 2010; Alvarado et al., 2011; Shephard et al., 2011; Adon et al., 2013; R'Honi et al., 2013; Whitburn et al., 2015, 2016a; Benedict et al., 2017; Warner et al., 2017). Excess NH_3 in the environment is of great concern since it is responsible for many environmental issues such as eutrophication of terrestrial and aquatic ecosystems, soil acidification and loss of plant diversity (Aneja et al., 2001; Erisman et al., 2007). As the dominant alkaline species in the atmosphere, NH_3 rapidly combines with acid gases such as sulfuric acid (H_2SO_4), nitric acid (HNO_3) and hydrochloric acid (HCl) resulting in the formation of secondary aerosols that are in turn impacting climate and human health (Bouwman et al., 1997; Aneja et al., 2001; Sutton et al., 2011; Behera et al., 2013; Lelieveld et al., 2015).

Until recently, most models of fire emissions were based on "bottom-up" approaches which rely on an estimation of the total burned biomass (BB, kg) combined with biome-specific emission factors (EFs), expressed as the mass of pollutant emitted per kilogram of BB (g kg^{-1} BB). Despite the numerous studies achieved in the past decades (e.g., Sinha et al., 2003; Yokelson et al., 2003; van der Werf et al., 2010; Wooster et al., 2011; Smith et al., 2014), the uncertainty on all parameters of these models remain large. This is especially true for EFs, which have typical uncertainty of the order of 20-30% for frequently measured species (e.g. CO, CO_2) and up to 100% for species such as NH_3 which are not so well monitored (Langmann et al., 2009; Akagi et al., 2011). An accurate determination of the EFs is challenging partly because of the existence of a within-biome spatial and seasonal variability (van Leeuwen and van der Werf, 2011; Yokelson et al., 2011; Meyer et al., 2012; Mebust and Cohen, 2013; van Leeuwen et al., 2013; Castellanos et al., 2014; Schreier et al., 2014a). This variability is attributed to differences in fuel type and burning conditions, the latter being itself controlled by climate, weather, moisture content, topography and fire practices (Andreae and Merlet, 2001; Korontzi et al., 2003; Yokelson et al., 2011; van Leeuwen and van der Werf, 2011; Castellanos et al., 2014). For nitrogen compounds, another main factor controlling the EFs is the Nitrogen content of the fuel (Andreae and Merlet, 2001; Jaffe and Wigder, 2012; Castellanos et al., 2014). Because it is generally not known to what extent EFs are based on a representative sample of a specific vegetation type (van Leeuwen and van der Werf, 2011; Castellanos et al., 2014), the spatial and temporal variability in the EFs is usually not taken into account in the bottom-up approaches where EFs are taken from compilations of airborne and local measurements or from small fires burned under laboratory conditions (e.g., Andreae and Merlet, 2001; Akagi et al., 2011).

With their excellent spatial and temporal coverage, hyperspectral sounders on board satellites, directly measuring tropospheric concentration of trace gases in the atmosphere, offer a unique opportunity to determine EFs more accurately and to capture their variability in time and space. Nowadays, the focus was principally on CO, nitrogen dioxide (NO_2) and aerosols (e.g., Pechony et al., 2013; Castellanos et al., 2014; Ichoku and Ellison, 2014; Mebust and Cohen, 2014; Schreier et al., 2014a, b). A recent study was also dedicated to formic acid (HCOOH) (Pommier et al., 2017). Until now, less attention has been given to NH_3 (Coheur et al., 2009; Alvarado et al., 2011; R'Honi et al., 2013; Luo et al., 2015). In this paper we derive biome-specific

NH₃ enhancement ratios relative to CO (ER_{NH₃/CO}, also known as normalized excess mixing ratios) and relate them to EFs (see section 2.2), over large tropical fire regions and long periods using the measurements of the Infrared Atmospheric Sounding Interferometer (IASI). The use of IASI is particularly suitable here because of its exceptional sampling (compared to other similar instruments, such as the Tropospheric Emission Spectrometer (TES) (Shephard et al., 2011)) and to our knowledge, it is the first time such a study focusing on biomass burning ERs is carried out at this scale for NH₃. Section 2 hereafter briefly describes the datasets used and introduces the methodology for calculating the enhancement ratios. It also motivates the selection of the regions studied. The results from our analyzes are presented and discussed in section 3, which is further divided in two main parts. The first part analyzes the variability of ER_{NH₃/CO} between and within the different biomes (an extensive comparison with ERs reported in the literature is also provided) while the second analyzes the interannual and seasonal evolution of ER_{NH₃/CO}. Summary and conclusion are given in section 4.

2 Dataset and Method

2.1 Instruments and data products

IASI is a nadir-looking high resolution Fourier Transform Spectrometer on board the polar orbiting sun-synchronous Metop (Meteorological Operational) satellites. The two first IASI sounders were launched in 2006 and 2012 (Metop-A and -B, respectively). A third instrument is scheduled for launch in 2018 and will ensure at least 18 years of consistent measurements (2006–2023). IASI covers the entire globe twice daily (09:30 and 21:30 local time when crossing the equator) with a relatively small elliptical footprint on the ground varying from 12 × 12 km (at nadir) up to 20 × 39 km (off nadir); depending on the viewing angle (Clerbaux et al., 2009). Its large and continuous spectral coverage of the thermal infrared band region (645–2760 cm⁻¹), its medium spectral resolution (0.5 cm⁻¹ apodized) and its low instrumental noise (~0.2 K at 950 cm⁻¹ and 280 K) make it an invaluable instrument for monitoring atmospheric composition (Clerbaux et al., 2009). CO is retrieved from IASI measurements using the FORLI (Fast Optimal Estimation Retrievals on Layers for IASI) software (Hurtmans et al., 2012) which has been extensively described in Hurtmans et al. (2012). The retrieval of NH₃ is based on a new and flexible retrieval algorithm, which relies on the calculation of a so-called Hyperspectral Range Index (HRI) and subsequent conversion to a NH₃ total column (molec.cm⁻²) using a neural network (Whitburn et al., 2016b). The retrieval also includes a full uncertainty analysis, performed by perturbing the input parameters (temperature profile, HRI, NH₃ a priori profile, etc.) of the neural network. We use in this paper the ANNI-NH₃-v2R-I version of the product which relies on ERA-Interim ECMWF meteorological input data, along with built-in surface temperature (Van Damme et al., 2017). For a detailed description of the NH₃ retrieval methods and parameters~~the parameters used in the ANNI-NH₃-v2R-I dataset~~, we refer the reader to ~~Whitburn et al. (2016b) for NH₃ and Hurtmans et al. (2012) for CO~~Whitburn et al. (2016b) and Van Damme et al. (2017). The validation of FORLI-CO profiles and columns have shown good agreement overall using in-situ, aircraft and satellite observations (Pommier et al., 2010; De Wachter et al., 2012; Kerzenmacher et al., 2012; George et al., 2015). For NH₃ columns, the validation has started but is more difficult considering the important spatial and temporal variability of NH₃ and the scarcity of correlative ground- and airplane-based measurements in many regions of the world (Van Damme et al., 2015). Two studies, based on a previous NH₃ re-

trieval algorithm also using the HRI but relying on a two-dimensional look-up tables for the conversion into a NH_3 total column (molec.cm^{-2}) (Van Damme et al., 2014), have shown fair agreements between IASI- NH_3 observations and other measurements (Van Damme et al., 2015; Dammers et al., 2016) (generally within the uncertainties of the IASI NH_3 retrieved columns), with differences of about 60–80% reported in Van Damme et al. (2015) and of 30% on average in Dammers et al. (2016).

- 5 This work makes use of 8 years (2008–2015) of daily global NH_3 and CO total columns (molec.cm^{-2}) from the measurements of IASI on board Metop-A. Only daytime satellite observations have been considered as they usually show a better sensitivity, especially to NH_3 . We also have assumed a similar sensitivity for IASI to NH_3 and CO in the lower layers of the atmosphere, which. This is not expected to introduce a significant bias in the studied regions due to a generally since it has been shown for both CO and NH_3 that the peak sensitivity was in the lower layers of the atmosphere in case of positive thermal contrast prevailing which is generally prevailing in the studied regions during daytime (Clarisse et al., 2010; Van Damme et al., 2014) (George et al., 2016b). A more important bias may result from the use of a unique vertical profile shape in the retrieval scheme of NH_3 total columns which is therefore not representative of the large variety of profiles observed above biomass burning plumes. Whitburn et al. (2016b) have calculated that the use of an alternative profile could affect the retrieved column up to 50%. This is important to keep in mind for the analyses presented next.
- 15 We also used in support to the selection of the studied regions and the NH_3 and CO columns, active fires detection data and nitrogen dioxide (NO_2) total columns (molec.cm^{-2}). Detected active fires are taken from the Global Monthly Fire Location Product (MCD14ML, Level 2, Collection 5) developed by the University of Maryland from the measurements of the MODerate resolution Imaging Spectroradiometer (MODIS) on board the NASA Terra and Aqua satellites (Justice et al., 2002; Giglio et al., 2006). Active fires are monitored at a resolution of $1 \times 1 \text{ km}^2$ with fires as small as 100 m^2 detected. NO_2 total columns are retrieved from the measurements of the Global Ozone Monitoring Experiment (GOME-2) also on board the Metop satellites and working in the UV-Vis spectral band region (Valks et al., 2011).

2.2 Enhancement ratios

- From the IASI NH_3 and CO total columns (molec.cm^{-2}), we have derived NH_3 enhancement ratios relative to CO ($\text{ER}_{\text{NH}_3/\text{CO}}$) defined as the ratio of the number of emitted molecules of NH_3 (here the NH_3 total column) over the emitted molecules of the reference species CO (here the CO total column) (Andreae et al., 1988; Lefer et al., 1994; Hobbs et al., 2003). The choice of CO as reference species is here particularly suitable as it is a dominant species emitted by fires and has a lifetime of several weeks in the free troposphere. One main advantage of the ERs compared to the EFs is that ERs calculation only requires simultaneous measurements of the studied (NH_3) and the reference species (CO) while EFs calculation requires fuel information that are not always available or completely reliable (Andreae and Merlet, 2001). In fire plumes, ERs can be estimated following (Goode et al., 2000; R'Honi et al., 2013):

$$\text{ER}_{\text{NH}_3/\text{CO}} = \frac{[\text{NH}_3]_{\text{smoke}} - [\text{NH}_3]_{\text{ambient}}}{[\text{CO}]_{\text{smoke}} - [\text{CO}]_{\text{ambient}}} \quad (1)$$

When a lot of measurements are available, which is often the case for IASI-derived measurements owing to its excellent spatial and temporal resolution, average $ER_{NH_3/CO}$ can be estimated from the slope of the linear regression of NH_3 versus CO (Andreae and Merlet, 2001; Coheur et al., 2009). The ERs can also be derived directly from the EFs by multiplying the ratio EF_{NH_3}/EF_{CO} with the ratio of the molar masses M_{CO}/M_{NH_3} (Andreae and Merlet, 2001). This will be used here to convert the reported EF values from ground-based and airborne studies into ERs in order to allow the comparison with our IASI-derived $ER_{NH_3/CO}$.

2.3 Selection of the areas and biomes and calculation of the $ER_{NH_3/CO}$

One of the key steps in this study is the selection of the areas of interest for the calculation of the $ER_{NH_3/CO}$. To be relevant, $ER_{NH_3/CO}$ need to be calculated for areas where fires are the dominant source of emissions of NH_3 and CO. The selection has been done on a pixel basis. We have first calculated the linear regressions, globally on a $1^\circ \times 1^\circ$ grid, between the monthly means of the couples NH_3 –CO total columns (molec.cm⁻²), NH_3 – NO_2 total columns, and NH_3 total columns–number of active fires (#fires). We have next selected the pixels for which a correlation coefficient (r) higher than 0.3 was found for the three couples of regression (NH_3 –CO, NH_3 – NO_2 and NH_3 –#fires). These are shown in Fig. 1 (colored pixels) and constitute the areas considered for the calculation of the $ER_{NH_3/CO}$. Pixels with a r higher than 0.3 for the considered couple but not for (at least) one of the two other couples are shown in gray. The idea behind this selection procedure is that a good correspondence between the monthly means of NH_3 , CO and NO_2 total columns provides an indication of a dominant contribution of the fires to their emissions, since biomass burning is indeed the only major common source of emissions of these three species. A significant positive correlation between the NH_3 total columns and the detected number of fires adds an additional argument in favor of the contribution of fires and ensures keeping only those areas that are close to the source of emissions, making the comparison with ground-based and airborne derived EFs and ERs easier. In general, the largest correlations are found between NH_3 and CO total columns (Fig. 1, top panel) with correlation coefficients ranging from about 0.6–0.7 up to 0.9 in Africa South of the Equator and Indonesia. The fact that these two species are measured simultaneously from IASI could contribute to this. For the two other couples (NH_3 – NO_2 and NH_3 –#fires), the correlation coefficients are in the range 0.3–0.8. Note that in general, significant positive correlations between NH_3 and NO_2 (Fig. 1, middle panel) are only found close to the source of emissions due to the relatively short lifetime of NO_2 (of a few hours, Schreier et al., 2014b). With a lifetime of typically 12–36 hours in the studied regions (Dentener and Crutzen, 1994; Aneja et al., 2001; Whitburn et al., 2015, 2016a), NH_3 is more likely to be transported over longer distances. This can be seen on the NH_3 – CO correlation map where positive correlations are also found over seas downwind of the source areas.

For each of the selected pixels, we have next calculated an $ER_{NH_3/CO}$ per year between 2008 and 2015 from the slope of the linear regression between NH_3 and CO retrieved columns (molec.cm⁻²). [The method considered here for the calculation of the regression line was the ordinary least square fit.](#) To take into account the NH_3 and CO columns most likely related to fire emissions, we have only considered IASI measurements located within 50 km from a fire. We have also included a quality filter on the NH_3 and CO measurements: only total columns with a relative error lower than 100% for NH_3 and 25% for CO were retained for the regression. Finally, as a post-filtering, we have only kept for the analysis the $ER_{NH_3/CO}$ for which the linear

regressions between NH_3 and CO columns show a correlation coefficient larger than 0.3 and for which we have more than 10 measurements. The impact of these pre- and post-filters on the calculated $\text{ER}_{\text{NH}_3/\text{CO}}$ are discussed in section 3.1. An example of a linear regression between NH_3 and CO for one of the selected pixels (Evergreen Broadleaf Forest (EBF) in Indonesia) is given in Fig. 2.

- 5 For this study, we focus on the four dominant biomes in the selected pixels. These were identified using the MODIS *Land Cover Type* product (MCD12Q1) with the 17-class International Geosphere-Biosphere Program classification (IGBP) (Friedl et al., 2010) (Fig. 3). The four selected classes are 1) the Evergreen Broadleaf Forests (EBF), 2) the Woody Savannas (WS), 3) the Savannas (S) and 4) the Crops together with the Crop and Natural Vegetation Mosaic (C+CNVM), here denoted C. Fig. 3 also shows the distribution of the mean yearly $\text{ER}_{\text{NH}_3/\text{CO}}$ averaged over the time period 2008–2015 for the selected pixels.
- 10 A first analysis of the distribution of the $\text{ER}_{\text{NH}_3/\text{CO}}$ reveals a variability between the four biomes, especially in Africa North of the Equator and in Central South America where a gradient is observed between EBF and WS and between EBF and S, respectively, with higher $\text{ER}_{\text{NH}_3/\text{CO}}$ found for EBF. A clear gradient is observed as well in Africa South of the Equator from the northwest to the southeast.

- The pixel-based $\text{ER}_{\text{NH}_3/\text{CO}}$ have next been grouped by biome to analyze their regional and temporal variability. In addition,
- 15 to facilitate the study of the spatial distribution of the $\text{ER}_{\text{NH}_3/\text{CO}}$, we have defined six main regions which include the majority of the pixels of interest (see Fig. 1). Two are in Africa, one North (AFR.NEQ.) and one South (AFR.SEQ.) of the Equator. One corresponds to the central part of South America (S.AM.). A second region in America (C.AM.) is located north to the S.AM. region and includes the region around the Gulf of Mexico, Central America, Colombia and Venezuela. The two last regions are in Asia; one is for South-East Asia (SE. ASIA) and the second is for Indonesia (INDO.).

20 3 Results and discussion

3.1 $\text{ER}_{\text{NH}_3/\text{CO}}$ spatial analysis

- We analyze here the spatial and biome variability in the $\text{ER}_{\text{NH}_3/\text{CO}}$. For each of the 4 biomes (EBF, WS, S, C) and each of the 6 regions a mean $\text{ER}_{\text{NH}_3/\text{CO}}$ was obtained by averaging all yearly pixel based $\text{ER}_{\text{NH}_3/\text{CO}}$ in the time period 2008 and 2015 (Fig. 4, solid error bars). Mean $\text{ER}_{\text{NH}_3/\text{CO}}$ for the 6 regions globally are shown as well (horizontal lines). Overall, the
- 25 highest mean $\text{ER}_{\text{NH}_3/\text{CO}}$ is found for EBF ($9.815.3 \times 10^{-3}$) while S and WS show mean $\text{ER}_{\text{NH}_3/\text{CO}}$ about ~~75-100~~40-50% lower with values of ~~5.99~~1 $\times 10^{-3}$ and ~~7.010.4~~10.4 $\times 10^{-3}$, respectively. The larger $\text{ER}_{\text{NH}_3/\text{CO}}$ for Evergreen Broadleaf Forests (EBF) compared to Woody Savannas (WS) and Savannas (S) is in agreement with previous studies (e.g. Andreae and Merlet, 2001; Akagi et al., 2011; Yokelson et al., 2011) and is mainly attributed to differences in fuel size and density: EBF, characterized by dense fuel, is indeed dominated by smoldering combustion, which emits more reduced or incompletely oxidized products
- 30 (among them NH_3 and CO) than grassland (van Leeuwen and van der Werf, 2011). One should note, however, that Kaiser et al. (2012) reported higher $\text{ER}_{\text{NH}_3/\text{CO}}$ for Savannas than for tropical forests. For Crop (C), the mean $\text{ER}_{\text{NH}_3/\text{CO}}$ (~~9.012.8~~12.8 $\times 10^{-3}$) is close to the $\text{ER}_{\text{NH}_3/\text{CO}}$ (EBF) but is more difficult to interpret because the biome probably includes different types of fuel. Figure 5, representing the cumulative frequency of the pixel-based yearly $\text{ER}_{\text{NH}_3/\text{CO}}$ per biome, also shows the biome-trends

in the $ER_{NH_3/CO}$. EBF ~~and C have~~ has for example about ~~35~~40% of the calculated $ER_{NH_3/CO}$ above ~~0.01~~0.015 while this value corresponds to only about ~~10–155–10~~% for S and WS. These differences in the $ER_{NH_3/CO}$ between biomes are, however, not necessarily found when looking at the average $ER_{NH_3/CO}$ at the region scale. For ~~Central America (C.A.M.) and~~ South-East Asia (SE.ASIA) in particular, the differences between $ER_{NH_3/CO}$ are low (of the order of 5–10%). For South America (S.A.M.),

5 $ER_{NH_3/CO}$ (EBF) is about ~~twice~~60% higher than $ER_{NH_3/CO}$ (S) but close to $ER_{NH_3/CO}$ (WS) (within 10%).

When comparing the $ER_{NH_3/CO}$ by biome between the six regions in Figure 4 (solid error bars), we find good agreements but also large differences, in line with what has already been reported by, for example, van Leeuwen and van der Werf (2011); van Leeuwen et al. (2013); Castellanos et al. (2014). Among the most noticeable differences, we find $ER_{NH_3/CO}$ (EBF) ~~up to a factor two higher~~ for the region of Africa North of the Equator (AFR.NEQ.) between 20% and 40% higher than for

10 the S.A.M., C.A.M. and SE.ASIA regions. Similarly, a large variability in the $ER_{NH_3/CO}$ is found for the WS and S biomes ranging between about ~~5~~7 $\times 10^{-3}$ for the region (s)-of AFR.SEQ. ~~(and S.A.M.) for WS (S) and and~~ 19 $\times 10^{-3}$ (~~14~~14 $\times 10^{-3}$ ~~for C~~) for S.A.M. (S.C.A.M.) for S (WS). A variability is also observed, but to a lesser extent, for WS (S). For the C biomeranging between 7.3, almost no variability is observed with $ER_{NH_3/CO}$ ranging between 11 $\times 10^{-3}$ for SE.ASIA and ~~10.5~~about 14 $\times 10^{-3}$ for C.A.M. ~~and AFR.NEQ.~~ Note that this intra-biome variability is also found within a given region, as observed in Fig. 3 and

15 as evidenced by the sometimes large standard deviation (std) associated with the mean $ER_{NH_3/CO}$ (e.g. EBF in the AFR.NEQ. region with a std ~~of about~~ higher than 0.01). As mentioned in section 1, these differences can be explained by changes in the fuel type (size and density) but also the climate, weather, topography, moisture and N content, and fire practices. In addition for EBF, different regional deforestation practices could also lead to variation in the $ER_{NH_3/CO}$ (van Leeuwen and van der Werf, 2011). It should finally be mentioned that for the AFR.NEQ. region, the measured NH_3 columns at the end of the fire

20 period probably originate from the combination of both biomass burning emissions and another source, possibly agriculture as suggested in Whitburn et al. (2015); this might therefore introduce a bias in the $ER_{NH_3/CO}$. Overall, these results clearly highlight the need for developing new regional-dependent EFs, in order to improve the representativeness of estimations from bottom-up inventories.

The comparison of the regional IASI-derived $ER_{NH_3/CO}$ (Fig. 4, solid error bars) with the values reported in the literature from ground-based or airborne studies (see Table 1) shows a ~~general overestimation of the latter (or an underestimation from IAS)~~good correspondence, especially for the ~~biomes EBF and C where a factor 1.5–4 difference is observed~~EBF and the S/WS biomes where $ER_{NH_3/CO}$ are generally within the range of values given in the literature. The only exception is for ~~Kaiser et al. (2012) for EBF for which a better correspondence (within 50% difference) is found. Note that the largest difference for EBF is with an~~ Yokelson et al. (2011) for EBF who measured $ER_{NH_3/CO}$ ~~reported for of about a factor 2–3 higher. The latter was however derived from~~ tropical dry forest ~~(Yokelson et al., 2011), but the latter and~~ is likely not representative for the complete Evergreen Broadleaf Forest class. ~~For S and WS, the agreement is much better (with a maximum underestimation of about 40%) except for Andreae and Merlet (2001), Bertschi et al. (2003) and Kaiser et al. (2012) reporting~~ $ER_{NH_3/CO}$ up to a factor 4 higher. Note that, in Bertschi et al. (2003) $ER_{NH_3/CO}$ are derived from smoldering logs for which higher values are logically expected. Note also that for Note that for WS, the $ER_{NH_3/CO}$ are compared here to values reported

30 for ~~savannas~~Savannas, which are usually included with S into the same biome. For croplands, values reported in the literature

are in contrast about 2–3 times higher than the one derived from IASI measurements. As mentioned before, the C biome probably include different type of fuels and results are therefore more difficult to interpret.

When looking at the mean $ER_{NH_3/CO}$ averaged over the six regions (Fig. 4, horizontal lines) for the four biomes, we find that the latter generally fall in the lower bound of the range given by the $ER_{NH_3/CO}$ reported in the literature. While an overestimation of the average $ER_{NH_3/CO}$ (or EF_{NH_3}) in the literature is possible, other reasons are likely to play a role. First, the differences with the IASI-derived $ER_{NH_3/CO}$ could also be (at least partly) explained by the consideration in our work of IASI measurements within 50 km of an active fire, while ground and airborne measurements are done in the direct vicinity of the fire. Second, another possible reason might lie in the difficulty for MODIS to detect smoldering fires, causing the IASI-derived $ER_{NH_3/CO}$ to reflect preferentially the flaming phase of the vegetation fires. Third, an accumulation of CO in the region during the fire period (due to its much longer lifetime compared to NH_3) might introduce a bias in the calculated $ER_{NH_3/CO}$. Finally, the differences with the reported $ER_{NH_3/CO}$ could also be due to the chosen methodology for the calculation of the $ER_{NH_3/CO}$. To verify this, we have recalculated mean biome-specific $ER_{NH_3/CO}$ for the six regions (not shown) by varying one by one the pre- and post-filters considered before (see section 2.3). We have performed four tests: 1) with a maximum distance of the NH_3 total column to a detected fire of 30 km and 2) 100 km (instead of 50 km), 3) with a maximum error on the NH_3 total column of 75% (against 100%) and 4) by filtering out the $ER_{NH_3/CO}$ for which the linear regressions between NH_3 and CO columns show a correlation coefficient (r) of the linear regression lower than 0.6 (instead of 0.3). We find a very limited impact of the distance to a fire and the error on the NH_3 column allowed, with differences of only about 3–8% (interestingly, an increase–decrease– of the tolerance on the maximum distance to a fire systematically slightly decrease–increase– the mean $ER_{NH_3/CO}$). In contrast, an increase to 0.6 of the threshold on the correlation coefficient introduces a large increase in the mean $ER_{NH_3/CO}$ of about 25–35% (and up to 40% for WS in the AFR.SEQ. region). 13–28%. Taking into account this increase, we find mean $ER_{NH_3/CO}$ closer (but still below) to the middle range of what is reported in the literature, especially for WS and S. The agreement would become even better if we consider in addition a possible bias due to the use of a non-representative constant NH_3 vertical profile considered for the retrieval of the NH_3 , as mentioned in section 2.1. Finally, as we mentioned in section 2.1, the retrieval of NH_3 could be biased by the use of a non-representative constant NH_3 vertical profile considered for the retrieval of the NH_3 , as mentioned in section 2.1. Note that despite the impact of the pre- and post-filters chosen, the analysis on the regional and inter-biome variability in the $ER_{NH_3/CO}$ remains valid.

At a regional level (all biomes combined), a comparison with the satellite-derived $ER_{NH_3/CO}$ based on TES measurements (Luo et al., 2015) shows again higher values of about 50% compared to an excellent agreement with our calculated $ER_{NH_3/CO}$ (Table 1). However, the calculation of $ER_{NH_3/CO}$ in Luo et al. (2015) are derived from seasonal averages in a $2^\circ \times 4^\circ$ grid and might therefore include other sources of emissions in the studied regions, such as agriculture and livestock. Luo et al. (2015) also derived $ER_{NH_3/CO}$ from simulations of the GEOS-Chem global chemical transport model. A good agreement is found between IASI and GEOS-Chem for the regions of AFR.NEQ. and S.AM. with $ER_{NH_3/CO}$ in the range of values calculated for North-Central Africa and South America. For South-Central Africa in contrast, Luo et al. (2015) reported $ER_{NH_3/CO}$ values of about 3–2 times higher compared to our AFR.SEQ. region.

3.2 $ER_{NH_3/CO}$ interannual and seasonal variability

In this second part, we focus our analysis on the temporal variability in the $ER_{NH_3/CO}$. Fig. 6 shows the mean $ER_{NH_3/CO}$ averaged by biome and by year (2008–2015). The solid line represents the 8-years average for each biome. We find an interannual variability in the mean $ER_{NH_3/CO}$ up to a factor two for the four studied biomes. ~~Interestingly, the~~ The minimum $ER_{NH_3/CO}$ is found in ~~2012 for all biomes. Similarly 2013 for the S and WS biomes while, for EBF, a minimum is observed in 2012.~~ Interestingly the highest mean $ER_{NH_3/CO}$ is observed in 2010 for all biomes (especially marked for EBF) except for ~~S-C~~ for which the maximum is found in ~~2008-2011~~ (despite an $ER_{NH_3/CO}$ for 2010 also above the 8 years average). When analyzing the variability in the yearly averaged $ER_{NH_3/CO}$ for each region separately (Fig. 7), we find that the high mean $ER_{NH_3/CO}$ of 2010 for EBF is exclusively carried by the AFR.NEQ. region with a mean $ER_{NH_3/CO}$ of ~~4035~~ $\times 10^{-3}$ (against ~~15~~ $\times 10^{-3}$ for the other years in the region). For the WS biome, the peak of 2010 is mainly due to the S.AM., AFR.NEQ. and SE.ASIA regions with an $ER_{NH_3/CO}$ about a factor 1.5–2.5 higher compared to the other years. This important variability in the $ER_{NH_3/CO}$ ~~are-is~~ probably due to differences in the burning conditions from one year to another. One possible reason to explain the high mean $ER_{NH_3/CO}$ for 2010 in the different regions is the El Niño Southern Oscillation (ENSO) event that occurred that year and that was responsible for severe droughts and increased fire activity in the studied regions (Whitburn et al., 2015). This is however probably not sufficient to explain the ~~3-2~~ times increase for EBF for 2010 in the AFR.NEQ. region but no clear evidence of other processes influencing the $ER_{NH_3/CO}$ were found for that year. Surprisingly, the same increase in the $ER_{NH_3/CO}$ are not observed for the year 2015, which was the strongest El Niño year since 1997 (Chisholm et al., 2016). For WS, high $ER_{NH_3/CO}$ are in addition observed for 2011 for South and Central America (S.AM. and C.AM.). However, this has small impact on the global yearly $ER_{NH_3/CO}$, which is mainly driven by the two regions in Africa (AFR.NEQ. and AFR.SEQ.), representing about ~~33%-and 4020% and 52%~~ of all the calculated $ER_{NH_3/CO}$ for WS, respectively (see Fig. 4). For the S biome, the yearly $ER_{NH_3/CO}$ is largely dominated by the AFR.SEQ. and the S.AM. (~~47%-and 3352% and 35%~~ of the $ER_{NH_3/CO}$, respectively). ~~Despite being not as pronounced on the global yearly $ER_{NH_3/CO}$ (Fig. 6) as for other biomes, we observe a peak~~ A peak is observed in the $ER_{NH_3/CO}$ for 2010 for the AFR.NEQ., the S.AM. and the C.AM. regions. Note that here the AFR.NEQ. and C.AM. regions also show high $ER_{NH_3/CO}$ for the year 2015 which tends to support the hypothesis of the influence of El Niño on the $ER_{NH_3/CO}$. Finally, the C biome is mainly driven by the AFR.NEQ. ~~region (60%-and the C.AM. regions (45% and 35% of all $ER_{NH_3/CO}$) showing as well the highest mean $ER_{NH_3/CO}$, respectively).~~ For the AFR.NEQ. region, a peak is observed in 2010 and ~~2015-2012~~ while for the C.AM. region, a maximum is reached in 2011 with an $ER_{NH_3/CO}$ about 50% higher compared to the other years. This important variability in the $ER_{NH_3/CO}$ in time and space highlights here again the importance to use dynamic EFs datasets in the fire emission inventories in order to better take into account the local fire conditions.

Finally we investigate the temporal variability in the $ER_{NH_3/CO}$ from a seasonal perspective. For this, we have calculated for each pixel selected in section 2.3 a separate $ER_{NH_3/CO}$ for the early and for the late fire season. The separation early/late fire season has been done by analyzing the daily time series of the number of fires between 2008 and 2015 for each region and biome studied (not shown). The results are shown in Fig. 4 (dashed error bars). In general, we do not find systematic difference in the $ER_{NH_3/CO}$ between the early and late fire season except for the AFR.NEQ. region, for which the late $ER_{NH_3/CO}$ are higher

by about ~~15-40~~20-40% for the four biomes. This is in agreement with the hypothesis made in section 3.1 of the presence of a secondary source of NH_3 (possibly agriculture) towards the end of the fire season. The same difference in the $\text{ER}_{\text{NH}_3/\text{CO}}$ was also observed by Luo et al. (2015) who found a 60% increase between the beginning and the end of the fire season for North-Central Africa. Note finally that the early and late fire season $\text{ER}_{\text{NH}_3/\text{CO}}$ are generally close to the corresponding yearly $\text{ER}_{\text{NH}_3/\text{CO}}$ (within 10-30%) which tend to support our methodology for the calculation of the ERs.

4 Conclusions

In this work, we have calculated biomass burning $\text{ER}_{\text{NH}_3/\text{CO}}$ over large tropical regions and a 8-year period of IASI satellite measurements for four different biomes, namely Evergreen Broadleaf Forest (EBF), Woody Savannas (WS), Savannas (S) and Cropland (C). Such a study had, to our knowledge, never been performed at this level (in time and space) for NH_3 .

Overall, the results have shown the great potential of IASI for calculating time and space dependent ERs. The $\text{ER}_{\text{NH}_3/\text{CO}}$ have been calculated on a pixel basis from the slope of the linear regression of NH_3 versus CO total columns (molec.cm^{-2}) retrieved from IASI measurements. On average, the biomes EBF and C showed $\text{ER}_{\text{NH}_3/\text{CO}}$ about ~~75-100~~40-50% higher than WS and S and this was attributed to differences in fuel size and density, affecting the fraction of smoldering combustion. The biome-specific $\text{ER}_{\text{NH}_3/\text{CO}}$ have next been grouped by regions and by year to analyze their spatial and temporal variability.

We found an important variability both in time and space for all situations but especially for WS showing a mean $\text{ER}_{\text{NH}_3/\text{CO}}$ about ~~50-100~~40-75% lower in Africa South of the Equator (AFR.SEQ.) than in the five other regions, possibly due to local differences in fuel type and burning conditions. Another interesting feature was the high mean $\text{ER}_{\text{NH}_3/\text{CO}}$ of ~~1723~~23 $\times 10^{-3}$ (and up to a factor two higher than for the other studied regions) calculated for Africa North of the Equator (AFR.NEQ.) for EBF. We have tentatively explained this high value by the presence of another source of emissions than biomass burning towards the end of the dry season. This was supported by our analysis of the seasonal dependence in the $\text{ER}_{\text{NH}_3/\text{CO}}$, showing $\text{ER}_{\text{NH}_3/\text{CO}}$ systematically higher for the late fire season in the AFR.NEQ. region (for the four biomes) than for the beginning of the fire period. The interannual variability in the $\text{ER}_{\text{NH}_3/\text{CO}}$ was also found important (up to a factor 2), with a peak for 2010 for each biome, possibly related to the severe droughts that have occurred that year in the studied regions consequently to an important El Niño event. The important variability of the $\text{ER}_{\text{NH}_3/\text{CO}}$ both in time and space clearly shows the need for developing dynamic datasets of EFs which better take into account the fuel type and fire conditions.

In comparison to the values reported in the literature, mainly from ground-based and airborne studies, ~~we found a general underestimation of the mean IASI-derived $\text{ER}_{\text{NH}_3/\text{CO}}$ of about 40% for Sand WS and up to a factor 1.5-4 for the EBF and C biomes. These differences for S, WS and EBF fell in the lower bound of the range given by the former. This~~ may be explained by various factors including 1) the parametrization (pre- and post-filtering of the data) considered for the calculation of the $\text{ER}_{\text{NH}_3/\text{CO}}$, 2) a bias towards the flaming phase due to the selection of IASI observations close to MODIS active fires (less sensitive to the smoldering phase) and 3) a possible accumulation of CO in the region during the fire season, introducing a low-bias in the IASI-derived $\text{ER}_{\text{NH}_3/\text{CO}}$. ~~The comparison with the recent work of Luo et al. (2015) who derived mean $\text{ER}_{\text{NH}_3/\text{CO}}$ from TES measurements tends to confirm a low-bias of the IASI-derived $\text{ER}_{\text{NH}_3/\text{CO}}$~~ Another possible explanation might lie in

the use of a unique vertical a priori profile in the retrieval scheme of NH_3 while biomass burning plumes exhibit a large variety of plume injection heights.

Data availability. The IASI FORLI CO and NH_3 Neural Network data used in this work are publicly available for all users through the French AERIS database (<http://iasi.aeris-data.fr/>)

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References

- Adon, M., Galy-Lacaux, C., Yoboué, V., Delon, C., Lacaux, J. P., Castera, P., Gardrat, E., Pienaar, J., Al Ourabi, H., Laouali, D., Diop, B., Sigha-Nkamdjou, L., Akpo, A., Tathy, J. P., Lavenu, F., and Mougin, E.: Long term measurements of sulfur dioxide, nitrogen dioxide, ammonia, nitric acid and ozone in Africa using passive samplers, *Atmospheric Chemistry and Physics*, 10, 7467–7487, doi:10.5194/acp-10-7467-2010, 2010.
- Adon, M., Galy-Lacaux, C., Delon, C., Yoboue, V., Solmon, F., and Kaptue Tchunte, A. T.: Dry deposition of nitrogen compounds (NO₂, HNO₃, NH₃), sulfur dioxide and ozone in west and central African ecosystems using the inferential method, *Atmospheric Chemistry and Physics*, 13, 11 351–11 374, doi:10.5194/acp-13-11351-2013, 2013.
- Akagi, S. K., Yokelson, R. J., Wiedinmyer, C., Alvarado, M. J., Reid, J. S., Karl, T., Crounse, J. D., and Wennberg, P. O.: Emission factors for open and domestic biomass burning for use in atmospheric models, *Atmospheric Chemistry and Physics*, 11, 4039–4072, doi:10.5194/acp-11-4039-2011, 2011.
- Alvarado, M. J., Cady-Pereira, K. E., Xiao, Y., Millet, D. B., and Payne, V. H.: Emission Ratios for Ammonia and Formic Acid and Observations of Peroxy Acetyl Nitrate (PAN) and Ethylene in Biomass Burning Smoke as Seen by the Tropospheric Emission Spectrometer (TES), *Atmosphere*, 2, 633–654, doi:10.3390/atmos2040633, 2011.
- Andreae, M. and Merlet, P.: Emission of trace gases and aerosols from biomass burning, *Global Biogeochemical Cycles*, 15, 955–966, 2001.
- Andreae, M., Browell, E., Garstang, M., Gregory, G. L., Harriss, R. C., Hill, G. F., Jacob, D., Pereira, M. C., Sachse, G. W., Setzer, A. W., and others: Biomass-burning emissions and associated haze layers over Amazonia, *Journal of Geophysical Research*, 93, 1509–1527, doi:10.1029/JD093iD02p01509, 1988.
- Aneja, V. P., Bunton, B., Walker, J. T., and Malik, B. P.: Measurement and analysis of atmospheric ammonia emissions from anaerobic lagoons, *Atmospheric Environment*, 35, 1949–1958, doi:10.1016/S1352-2310(00)00547-1, 2001.
- Bauduin, S., Clarisse, L., Theunissen, M., George, M., Hurtmans, D., Clerbaux, C., and Coheur, P. F.: IASI's sensitivity to near-surface carbon monoxide (CO): Theoretical analyses and retrievals on test cases., *Journal of Quantitative Spectroscopy & Radiative Transfer*, 189, 428–440, doi:10.1016/j.jqsrt.2016.12.022, 2016.
- Behera, S. N., Sharma, M., Aneja, V. P., and Balasubramanian, R.: Ammonia in the atmosphere: a review on emission sources, atmospheric chemistry and deposition on terrestrial bodies., *Environ. Sci. Pollut. Res. Int.*, 20(11), 8092–131, doi:10.1007/s11356-013-2051-9, 2013.
- Benedict, K. B., Prenni, A. J., Carrico, C. M., Sullivan, A. P., Schichtel, B. A., and Collett, J. L.: Enhanced concentrations of reactive nitrogen species in wildfire smoke, *Atmospheric Environment*, 148, 8–15, doi:10.1016/j.atmosenv.2016.10.030, 2017.
- Bertschi, I., Yokelson, R. J., Ward, D. E., Babbitt, R. E., Susott, R. A., Goode, J. G., and Hao, W. M.: Trace gas and particle emissions from fires in large diameter and belowground biomass fuels, *Journal of Geophysical Research*, 108, 8472, doi:10.1029/2002JD002100, 2003.
- Bouwman, A., Lee, D., Asman, W., Dentener, F., Van Der Hoek, K., and Olivier, J.: A global high-resolution emission inventory for ammonia, *Global Biogeochemical Cycles*, 11, 561–587, doi:10.1029/97GB02266, 1997.
- Bowman, D. M., Balch, Jennifer K. and Artaxo, P., Bond, W. J., Carlson, J. M., Cochrane, M. A., D'Antonio, C. M., DeFries, R. S., Doyle, J. C., Harrison, S. P., and Others: Fire in the Earth system, *Science*, 324, 481–484, doi:10.1126/science.1163886, 2009.
- Castellanos, P., Boersma, K. F., and van der Werf, G. R.: Satellite observations indicate substantial spatiotemporal variability in biomass burning NO_x emission factors for South America, *Atmospheric Chemistry and Physics*, 14, 3929–3943, doi:10.5194/acp-14-3929-2014, 2014.

- Chisholm, R. A., Wijedasa, L. S., and Swinfield, T.: The need for long-term remedies for Indonesia's forest fires: Letter, *Conservation Biology*, 30, 5–6, doi:10.1111/cobi.12662, 2016.
- Christian, T. J., Yokelson, R. J., Carvalho, J. a. A., Griffith, D. W. T., Alvarado, E. C., Santos, J. C., Neto, T. G. S., Veras, C. A. G., and Hao, W. M.: The tropical forest and fire emissions experiment: Trace gases emitted by smoldering logs and dung from deforestation and pasture fires in Brazil, *Journal of Geophysical Research*, 112, doi:10.1029/2006JD008147, 2007.
- Clarisse, L., Shephard, M., Dentener, F., Hurtmans, D., Cady-Pereira, K., Karagulian, F., Van Damme, M., Clerbaux, C., and Coheur, P.-F.: Satellite monitoring of ammonia: A case study of the San Joaquin Valley, *Journal of Geophysical Research*, 115, D13 302, doi:10.1029/2009JD013291, 2010.
- Clerbaux, C., Boynard, A., Clarisse, L., George, M., Hadji-Lazaro, J., Herbin, H., Hurtmans, D., Pommier, M., Razavi, A., Turquety, S., Wespes, C., and Coheur, P.-F.: Monitoring of atmospheric composition using the thermal infrared IASI/MetOp sounder, *Atmospheric Chemistry and Physics*, 9, 6041–6054, doi:10.5194/acp-9-6041-2009, 2009.
- Coheur, P.-F., Clarisse, L., Turquety, S., Hurtmans, D., and Clerbaux, C.: IASI measurements of reactive trace species in biomass burning plumes, *Atmospheric Chemistry and Physics*, 9, 5655–5667, doi:10.5194/acp-9-5655-2009, 2009.
- Crutzen, P. and Andreae, M.: Biomass burning in the tropics: impact on atmospheric chemistry and biogeochemical cycles., *Science*, 250, 1669–1678, doi:10.1126/science.250.4988.1669, 1990.
- Dammers, E., Palm, M., Van Damme, M., Vigouroux, C., Smale, D., Conway, S., Toon, G. C., Jones, N., Nussbaumer, E., Warneke, T., Petri, C., Clarisse, L., Clerbaux, C., Hermans, C., Lutsch, E., Strong, K., Hannigan, J. W., Nakajima, H., Morino, I., Herrera, B., Stremme, W., Grutter, M., Schaap, M., Wichink Kruit, R. J., Notholt, J., Coheur, P.-F., and Erisman, J. W.: An evaluation of IASI-NH₃ with ground-based Fourier transform infrared spectroscopy measurements, *Atmospheric Chemistry and Physics*, 16, 10 351–10 368, doi:10.5194/acp-16-10351-2016, 2016.
- De Wachter, E., Barret, B., Le Flochmoën, E., Pavelin, E., Matricardi, M., Clerbaux, C., Hadji-Lazaro, J., George, M., Hurtmans, D., Coheur, P.-F., Nedelec, P., and Cammas, J. P.: Retrieval of MetOp-A/IASI CO profiles and validation with MOZAIC data, *Atmospheric Measurement Techniques*, 5, 2843–2857, doi:10.5194/amt-5-2843-2012, 2012.
- Dentener, F. and Crutzen, P.: A Three-Dimensional Model of the Global Ammonia Cycle, *Journal of Atmospheric Chemistry*, 19, 331–369, 1994.
- Erisman, J., Bleeker, A., Galloway, J., and Sutton, M.: Reduced nitrogen in ecology and the environment, *Environmental Pollution*, 150, 140–149, doi:10.1016/j.envpol.2007.06.033, 2007.
- Friedl, M. A., Sulla-Menashe, D., Tan, B., Schneider, A., Ramankutty, N., Sibley, A., and Huang, X.: MODIS Collection 5 global land cover: Algorithm refinements and characterization of new datasets, *Remote Sensing of Environment*, 114, 168–182, doi:10.1016/j.rse.2009.08.016, 2010.
- Galloway, J., Dentener, F., Capone, D., Boyer, E., Howarth, R., Seitzinger, S., Asner, G., Cleveland, C., Green, P., Holland, E., Karl, D., Michaels, A., Porter, J., Townsend, A., and Vörösmarty, C.: Nitrogen cycles: past, present and future, *Biogeochemistry*, 70, 153–226, 2004.
- George, M., Clerbaux, C., Hurtmans, D., Turquety, S., Coheur, P.-F., Pommier, M., Hadji-Lazaro, J., Edwards, D. P., Worden, H., Luo, M., and others: Carbon monoxide distributions from the IASI/METOP mission: evaluation with other space-borne remote sensors, *Atmospheric chemistry and physics*, 9, 8317–8330, <https://hal.archives-ouvertes.fr/hal-00429693/>, 2009.
- George, M., Clerbaux, C., Bouarar, I., Coheur, P.-F., Deeter, M. N., Edwards, D. P., Francis, G., Gille, J. C., Hadji-Lazaro, J., Hurtmans, D., Inness, A., Mao, D., and Worden, H. M.: An examination of the long-term CO records from MOPITT and IASI: comparison of retrieval

- methodology, *Atmospheric Measurement Techniques*, 8, 4313–4328, doi:10.5194/amt-8-4313-2015, <http://www.atmos-meas-tech.net/8/4313/2015/>, 2015.
- Giglio, L., Van der Werf, G. R., Randerson, J. T., Collatz, G. J., and Kasibhatla, P.: Global estimation of burned area using MODIS active fire observations, *Atmospheric Chemistry and Physics*, 6, 957–974, doi:10.5194/acp-6-957-2006, 2006.
- 5 Goode, J. G., Yokelson, R. J., Ward, D. E., Susott, R. A., Babbitt, R. E., Davies, M. A., and Hao, W. M.: Measurement of excess O₃, CO₂, CO, CH₄, C₂H₄, C₂H₂, HCN, NO, NH₃, HCOOH, CH₃COOH, HCHO, and CH₃OH in 1997 Alaskan biomass burning plumes by airborne Fourier transform infrared spectroscopy (AFTIR), *Journal of Geophysical Research*, 105, 147–166, doi:10.1029/2000JD900287, 2000.
- Hobbs, P. V., Sinha, P., Yokelson, R. J., Christian, T. J., Blake, D. R., Gao, S., Kirchstetter, T. W., Novakov, T., and Pilewskie, P.: Evolution of gases and particles from a savanna fire in South Africa, *Journal of Geophysical Research*, 108, doi:10.1029/2002JD002352, 2003.
- 10 Hurtmans, D., Coheur, P.-F., Wespes, C., Clarisse, L., Scharf, O., O., Clerbaux, C., Hadji-Lazaro, J., George, M., and Turquety, S.: FORLI radiative transfer and retrieval code for IASI, *Journal of Quantitative Spectroscopy and Radiative Transfer*, 113, 1391–1408, doi:10.1016/j.jqsrt.2012.02.036, 2012.
- Ichoku, C. and Ellison, L.: Global top-down smoke-aerosol emissions estimation using satellite fire radiative power measurements, *Atmospheric Chemistry and Physics*, 14, 6643–6667, doi:10.5194/acp-14-6643-2014, 2014.
- 15 Jaffe, D. A. and Wigder, N. L.: Ozone production from wildfires: A critical review, *Atmospheric Environment*, 51, 1–10, doi:10.1016/j.atmosenv.2011.11.063, 2012.
- Justice, C. O., Giglio, L., Korontzi, S., Owens, J., Morisette, J., Roy, D., Descloitres, J., Alleaume, S., Petitcolin, F., and Kaufman, Y.: The MODIS fire products, *Remote Sensing of the Environment*, 83, 244–262, 2002.
- Kaiser, J. W., Heil, A., Andreae, M. O., Benedetti, A., Chubarova, N., Jones, L., Morcrette, J.-J., Razinger, M., Schultz, M. G., Suttie, M.,
20 and van der Werf, G. R.: Biomass burning emissions estimated with a global fire assimilation system based on observed fire radiative power, *Biogeosciences*, 9, 527–554, doi:10.5194/bg-9-527-2012, 2012.
- Kerzenmacher, T., Dils, B., Kumps, N., Blumenstock, T., Clerbaux, C., Coheur, P.-F., Demoulin, P., García, O., George, M., Griffith, D. W. T., Hase, F., Hadji-Lazaro, J., Hurtmans, D., Jones, N., Mahieu, E., Notholt, J., Paton-Walsh, C., Raffalski, U., Ridder, T., Schneider, M., Servais, C., and De Mazière, M.: Validation of IASI FORLI carbon monoxide retrievals using FTIR data from NDACC, *Atmospheric
25 Measurement Techniques*, 5, 2751–2761, doi:10.5194/amt-5-2751-2012, 2012.
- Korontzi, S., Ward, D. E., Susott, R. A., Yokelson, R. J., Justice, C. O., Hobbs, P. V., Smithwick, E. A. H., and Hao, W. M.: Seasonal variation and ecosystem dependence of emission factors for selected trace gases and PM_{2.5} for southern African savanna fires, *Journal of Geophysical Research: Atmospheres*, 108, doi:10.1029/2003JD003730, 2003.
- Langmann, B., Duncan, B., Textor, C., Trentmann, J., and van der Werf, G. R.: Vegetation fire emissions and their impact on air pollution
30 and climate, *Atmospheric Environment*, 43, 107–116, doi:10.1016/j.atmosenv.2008.09.047, 2009.
- Lefer, B., Talbot, R., Harriss, R., Bradshaw, J., Sandholm, S., Olson, J., Sachse, G. W., Collins, J., Shipham, M., Blake, D., Klemm, K., Klemm, O., Gorzelska, K., and Barrick, J.: Enhancement of acidic gases in biomass burning impacted air masses over Canada, *Journal of Geophysical Research*, 99, 1721–1737, doi:10.1029/93JD02091, 1994.
- Lelieveld, J., Evans, J. S., Fnais, M., Giannadaki, D., and Pozzer, A.: The contribution of outdoor air pollution sources to premature mortality
35 on a global scale, *Nature*, 525, 367–371, doi:10.1038/nature15371, 2015.
- Luo, M., Shephard, M. W., Cady-Pereira, K. E., Henze, D. K., Zhu, L., Bash, J. O., Pinder, R. W., Capps, S. L., Walker, J. T., and Jones, M. R.: Satellite observations of tropospheric ammonia and carbon monoxide: Global distributions, regional correlations and comparisons to model simulations, *Atmospheric Environment*, 106, 262–277, doi:10.1016/j.atmosenv.2015.02.007, 2015.

- Mebust, A. K. and Cohen, R. C.: Observations of a seasonal cycle in NO_x emissions from fires in African woody savannas, *Geophysical Research Letters*, 40, 1451–1455, doi:10.1002/grl.50343, 2013.
- Mebust, A. K. and Cohen, R. C.: Space-based observations of fire NO_x emission coefficients: a global biome-scale comparison, *Atmospheric Chemistry and Physics*, 14, 2509–2524, doi:10.5194/acp-14-2509-2014, 2014.
- 5 Meyer, C. P., Cook, G. D., Reisen, F., Smith, T. E. L., Tattaris, M., Russell-Smith, J., Maier, S. W., Yates, C. P., and Wooster, M. J.: Direct measurements of the seasonality of emission factors from savanna fires in northern Australia, *Journal of Geophysical Research: Atmospheres*, 117, 2156–2202, doi:10.1029/2012JD017671, 2012.
- Pechony, O., Shindell, D. T., and Faluvegi, G.: Direct top-down estimates of biomass burning CO emissions using TES and MOPITT versus bottom-up GFED inventory, *Journal of Geophysical Research: Atmospheres*, 118, 8054–8066, doi:10.1002/jgrd.50624, 2013.
- 10 Pommier, M., Law, K. S., Clerbaux, C., Turquety, S., Hurtmans, D., Hadji-Lazaro, J., Coheur, P.-F., Schlager, H., Ancellet, G., Paris, J.-D., Nédélec, P., Diskin, G. S., Podolske, J. R., Holloway, J. S., and Bernath, P.: IASI carbon monoxide validation over the Arctic during POLARCAT spring and summer campaigns, *Atmospheric Chemistry and Physics*, 10, 10 655–10 678, doi:10.5194/acp-10-10655-2010, 2010.
- Pommier, M., Clerbaux, C., and Coheur, P.-F.: Possibility for an infrared sounder as IASI to document the HCOOH chemistry in biomass burning plumes, *Atmospheric Chemistry and Physics Discussions*, pp. 1–22, doi:10.5194/acp-2017-126, 2017.
- 15 R'Honi, Y., Clarisse, L., Clerbaux, C., Hurtmans, D., Dufлот, V., Turquety, S., Ngadi, Y., and Coheur, P.-F.: Exceptional emissions of NH₃ and HCOOH in the 2010 Russian wildfires, *Atmospheric Chemistry and Physics*, 13, 4171–4181, doi:10.5194/acp-13-4171-2013, 2013.
- Schreier, S., Richter, A., Schepaschenko, D., Shvidenko, A., Hilboll, A., and Burrows, J.: Differences in satellite-derived NO_x emission factors between Eurasian and North American boreal forest fires, *Atmospheric Environment*, 121, 55–65, doi:10.1016/j.atmosenv.2014.08.071, 2014a.
- 20 Schreier, S. F., Richter, A., Kaiser, J. W., and Burrows, J. P.: The empirical relationship between satellite-derived tropospheric NO_x and fire radiative power and possible implications for fire emission rates of NO_x, *Atmospheric Chemistry and Physics*, 14, 2447–2466, doi:10.5194/acp-14-2447-2014, 2014b.
- Shephard, M. W., Cady-Pereira, K. E., Luo, M., Henze, D. K., Pinder, R. W., Walker, J. T., Rinsland, C. P., Bash, J. O., Zhu, L., Payne, V. H., and Clarisse, L.: TES ammonia retrieval strategy and global observations of the spatial and seasonal variability of ammonia, *Atmospheric Chemistry and Physics*, 11, 10 743–10 763, doi:10.5194/acp-11-10743-2011, 2011.
- 25 Sinha, P., Hobbs, P., Yokelson, R. J., Bertschi, I., Blake, D., Simpson, I., Gao, S., Kirchstetter, T., and Novakov, T.: Emissions of trace gases and particles from savanna fires in southern Africa, *Journal of Geophysical Research*, 108, doi:10.1029/2002JD002325, 2003.
- Smith, T. E. L., Paton-Walsh, C., Meyer, C. P., Cook, G. D., Maier, S. W., Russell-Smith, J., Wooster, M. J., and Yates, C. P.: New emission factors for Australian vegetation fires measured using open-path Fourier transform infrared spectroscopy – Part 2: Australian tropical savanna fires, *Atmospheric Chemistry and Physics*, 14, 11 335–11 352, doi:10.5194/acp-14-11335-2014, 2014.
- 30 Sutton, M., Oenema, O., Erisman, J., Leip, J., van Grinsven, H., and Winiwarter, W.: Too much of a good thing, *Nature*, 472, 159–161, 2011.
- Valks, P., Pinardi, G., Richter, A., Lambert, J.-C., Hao, N., Loyola, D., Van Roozendael, M., and Emmadi, S.: Operational total and tropospheric NO₂ column retrieval for GOME-2, *Atmospheric Measurement Techniques*, 4, 1491–1514, doi:10.5194/amt-4-1491-2011, 2011.
- 35 Van Damme, M., Clarisse, L., Heald, C. L., Hurtmans, D., Ngadi, Y., Clerbaux, C., Dolman, A. J., Erisman, J. W., and Coheur, P. F.: Global distributions, time series and error characterization of atmospheric ammonia (NH₃) from IASI satellite observations, *Atmospheric chemistry and physics*, 14, 2905–2922, doi:10.5194/acp-14-2905-2014, 2014.

- Van Damme, M., Clarisse, L., Dammers, E., Liu, X., Nowak, J. B., Clerbaux, C., Flechard, C. R., Galy-Lacaux, C., Xu, W., Neuman, J. A., Tang, Y. S., Sutton, M. A., Erisman, J. W., and Coheur, P. F.: Towards validation of ammonia (NH_3) measurements from the IASI satellite, *Atmospheric Measurement Techniques*, 8, 1575–1591, doi:10.5194/amt-8-1575-2015, 2015.
- Van Damme, M., Whitburn, S., Clarisse, L., Clerbaux, C., Hurtmans, D., and Coheur, P.-F.: Version 2 of the IASI NH_3 neural network retrieval algorithm; near-real time and reanalysed datasets, *Atmos. Meas. Tech. Discuss.*, in review, doi:10.5194/amt-2017-239, 2017.
- van der Werf, G. R., Randerson, J. T., Giglio, L., Collatz, G. J., Giglio, L., Kasibhatla, P., Arekkano, A., Olsen, S., and Kasischke, E.: Continental-Scale Partitioning of Fire Emissions During the 1997 to 2001 El Niño/La Niña Period, *Science*, 303, 73–76, doi:10.1126/science.1090753, 2004.
- van der Werf, G. R., Randerson, J. T., Giglio, L., Collatz, G. J., Mu, M., Kasibhatla, P., Morton, D. C., DeFries, R., Jin, Y., and van Leeuwen, T. T.: Global fire emissions and the contribution of deforestation, savanna, forest, agricultural, and peat fires (1997–2009), *Atmospheric Chemistry and Physics*, 10, 11 707–11 735, doi:10.5194/acp-10-11707-2010, 2010.
- van Leeuwen, T. T. and van der Werf, G. R.: Spatial and temporal variability in the ratio of trace gases emitted from biomass burning, *Atmospheric Chemistry and Physics*, 11, 3611–3629, doi:10.5194/acp-11-3611-2011, 2011.
- van Leeuwen, T. T., Peters, W., Krol, M. C., and van der Werf, G. R.: Dynamic biomass burning emission factors and their impact on atmospheric CO mixing ratios, *Journal of Geophysical Research: Atmospheres*, 118, 6797–6815, doi:10.1002/jgrd.50478, 2013.
- Warner, J. X., Dickerson, R. R., Wei, Z., Strow, L. L., Wang, Y., and Liang, Q.: Increased atmospheric ammonia over the world’s major agricultural areas detected from space, *Geophysical Research Letters*, 44, doi:10.1002/2016gl072305, 2017.
- Whitburn, S., Van Damme, M., Kaiser, J., van der Werf, G., Turquety, S., Hurtmans, D., Clarisse, L., Clerbaux, C., and Coheur, P.-F.: Ammonia emissions in tropical biomass burning regions: Comparison between satellite-derived emissions and bottom-up fire inventories, *Atmospheric Environment*, doi:10.1016/j.atmosenv.2015.03.015, 2015.
- Whitburn, S., Van Damme, M., Clarisse, L., Turquety, S., Clerbaux, C., and Coheur, P.-F.: Doubling of annual ammonia emissions from the peat fires in Indonesia during the 2015 El Niño, *Geophysical Research Letters*, 43, 11,007–11,014, doi:10.1002/2016GL070620, 2016a.
- Whitburn, S., Van Damme, M., Clarisse, L., Bauduin, S., Heald, C., Hadji-Lazaro, J., Hurtmans, D., Zondlo, M., Clerbaux, C., and Coheur, P.-F.: A flexible and robust neural network IASI- NH_3 retrieval algorithm, *Journal of Geophysical Research: Atmospheres*, 121, doi:10.1002/2016JD024828, 2016b.
- Wooster, M. J., Freeborn, P. H., Archibald, S., Oppenheimer, C., Roberts, G. J., Smith, T. E. L., Govender, N., Burton, M., and Palumbo, I.: Field determination of biomass burning emission ratios and factors via open-path FTIR spectroscopy and fire radiative power assessment: headfire, backfire and residual smouldering combustion in African savannahs, *Atmospheric Chemistry and Physics*, 11, 11 591–11 615, doi:10.5194/acp-11-11591-2011, 2011.
- Yokelson, R. J., Bertschi, I. T., Christian, T. J., Hobbs, P. V., Ward, D. E., and Hao, W. M.: Trace gas measurements in nascent, aged, and cloud-processed smoke from African savanna fires by airborne Fourier transform infrared spectroscopy (AFTIR), *Journal of Geophysical Research*, 108, doi:10.1029/2002JD002322, 2003.
- Yokelson, R. J., Burling, I. R., Urbanski, S. P., Atlas, E. L., Adachi, K., Buseck, P. R., Wiedinmyer, C., Akagi, S. K., Toohey, D. W., and Wold, C. E.: Trace gas and particle emissions from open biomass burning in Mexico, *Atmospheric Chemistry and Physics*, 11, 6787–6808, doi:10.5194/acp-11-6787-2011, 2011.

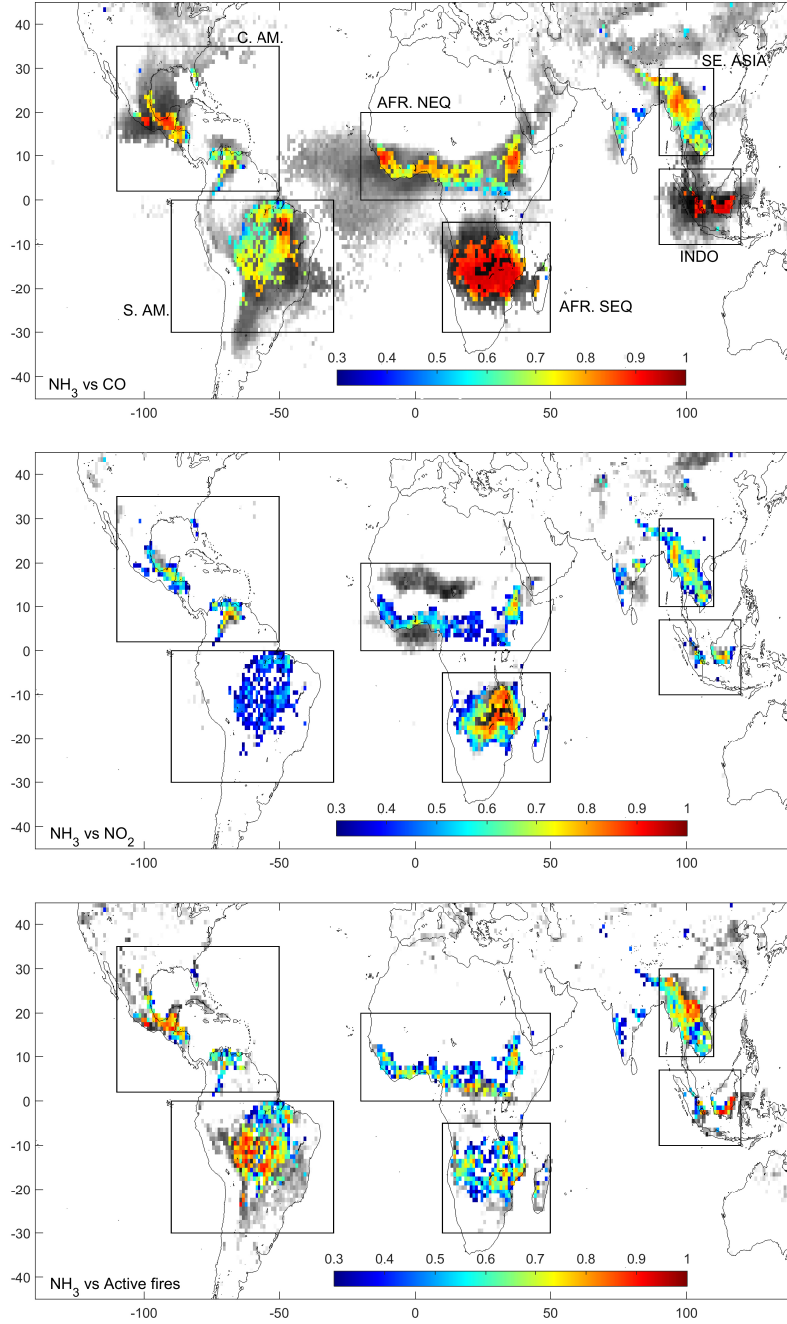


Figure 1. Correlation coefficients (r) of the linear regression of the monthly mean NH_3 total columns (molec.cm^{-2}) versus (a) CO total columns (molec.cm^{-2}) (top), (b) NO_2 total columns (molec.cm^{-2}) (middle) and (c) the number of active fires (bottom) from 2008 to 2015 in $1^\circ \times 1^\circ$ cells. Only pixels with a correlation coefficient r higher than 0.3 are shown. Pixels with $r > 0.3$ for the three couples of regression are shown in color. Pixels with $r > 0.3$ for the considered couple but not for (at least) one of the two other couples are shown in gray. The six regions selected for the study (C.AM., S.AM., AFR.NEQ., AFR.SEQ., SE.ASIA, INDO.) are highlighted by black rectangles.

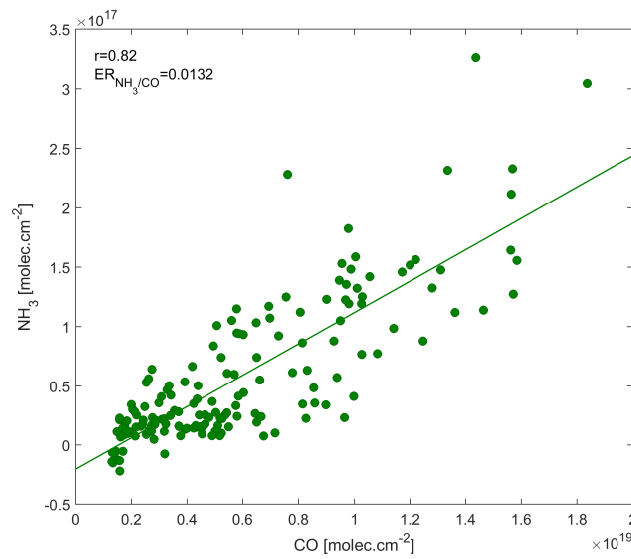


Figure 2. Example of linear regression between NH₃ and CO total columns (molec.cm⁻²) for 2015 for one pixel of the selected grid-box, corresponding to the Evergreen Broadleaf Forest (EBF) biome in Indonesia (latitude=-3 deg; longitude=113 deg). The correlation coefficient (r) and the ER_{NH₃/CO} (slope of the linear regression) are given as well as inset. The ordinary least square fit has been chosen here for the calculation of the regression line.

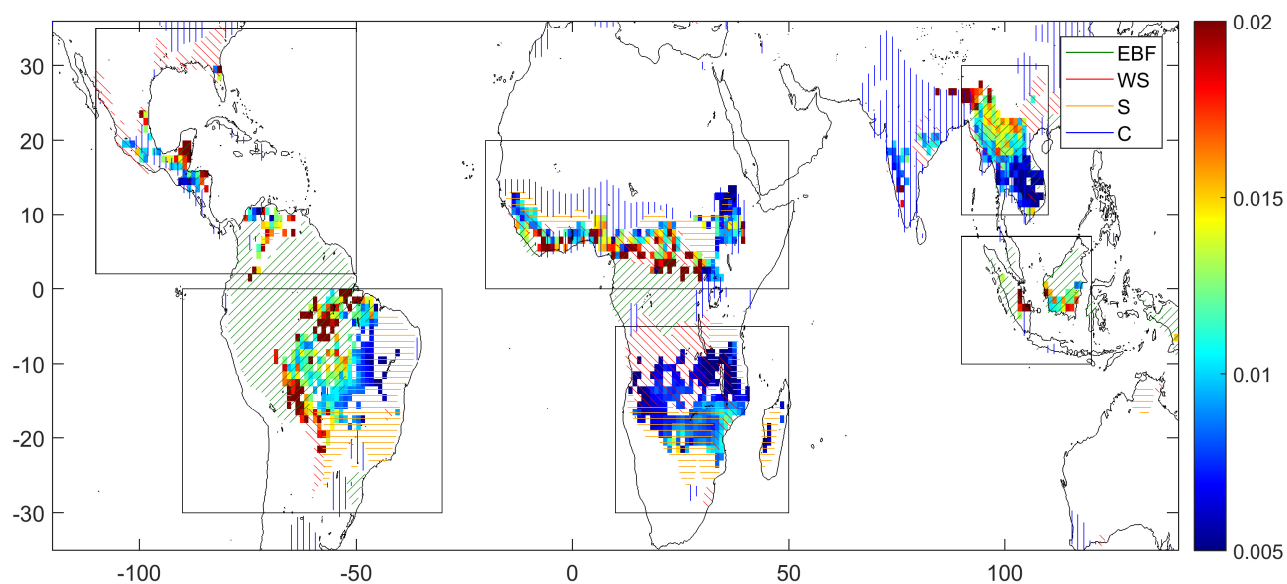


Figure 3. Mean yearly $ER_{NH_3/CO}$ averaged over the time period 2008 and 2015 for the selected pixels. The four main biomes studied are represented by the hatched lines: Savannas (S), Woody Savannas (WS), Evergreen Broadleaf Forest (EBF) and Crop together with the Crop and Natural Vegetation Mosaic (C+CNVM) here called as C.

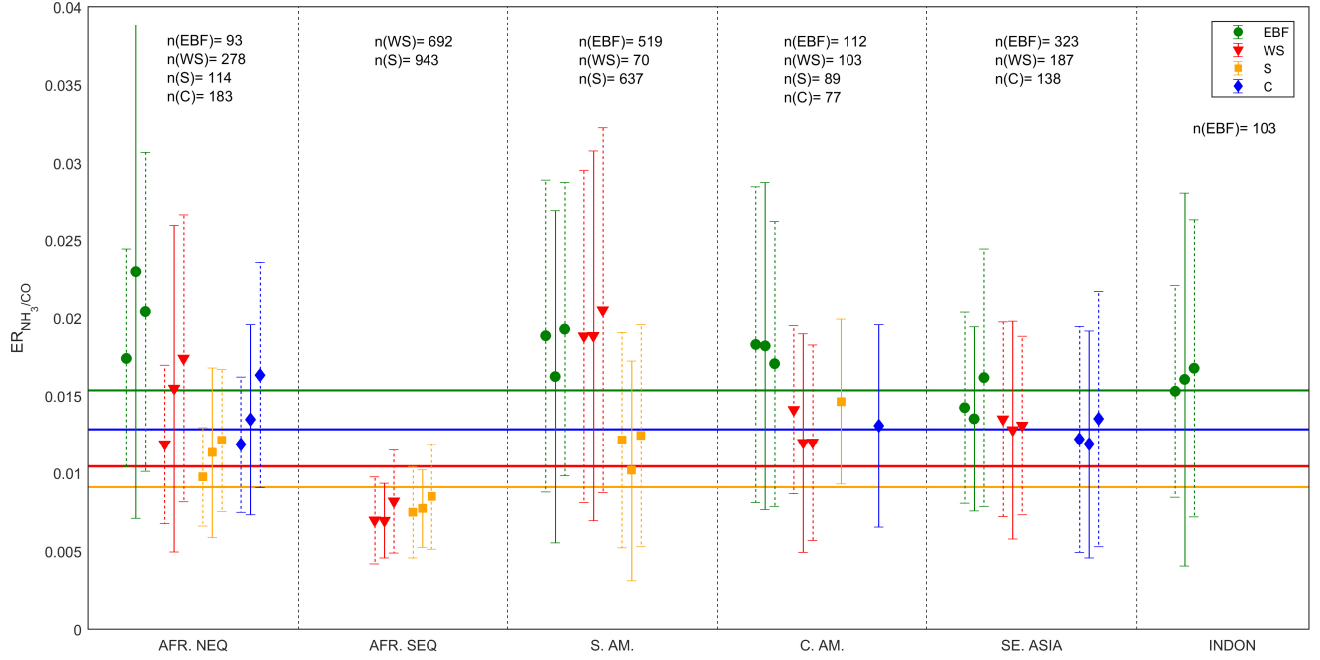


Figure 4. Mean $ER_{NH_3/CO}$ averaged for the six regions and four biomes from the yearly $ER_{NH_3/CO}$ (solid error bars) and from the early- and late fire season $ER_{NH_3/CO}$ (left and right dashed error bars, respectively) calculated between 2008 and 2015 for the pixels selected in section 2.3. The error bar is the 1-sigma standard deviation around the mean. The mean yearly $ER_{NH_3/CO}$ for each biome averaged globally for the six regions are indicated by the horizontal lines. $n(x)$ (with x the biome) corresponds to the number of $ER_{NH_3/CO}$ averaged for each biome and region. Different symbols and colors are used for the different biomes. For the S and C biomes in the C.AM. region, no seasonal $ER_{NH_3/CO}$ are shown because of the lack of measurements.

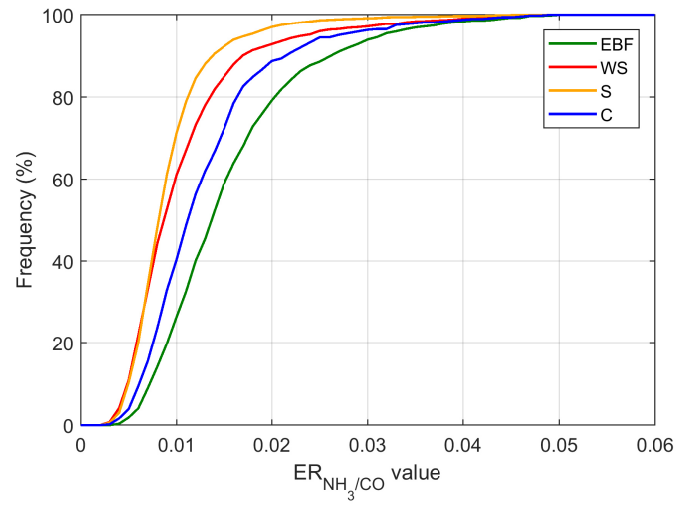


Figure 5. Cumulative curve of the yearly $ER_{NH_3/CO}$ calculated between 2008 and 2015 for the pixels selected in section 2.3 separated by biome.

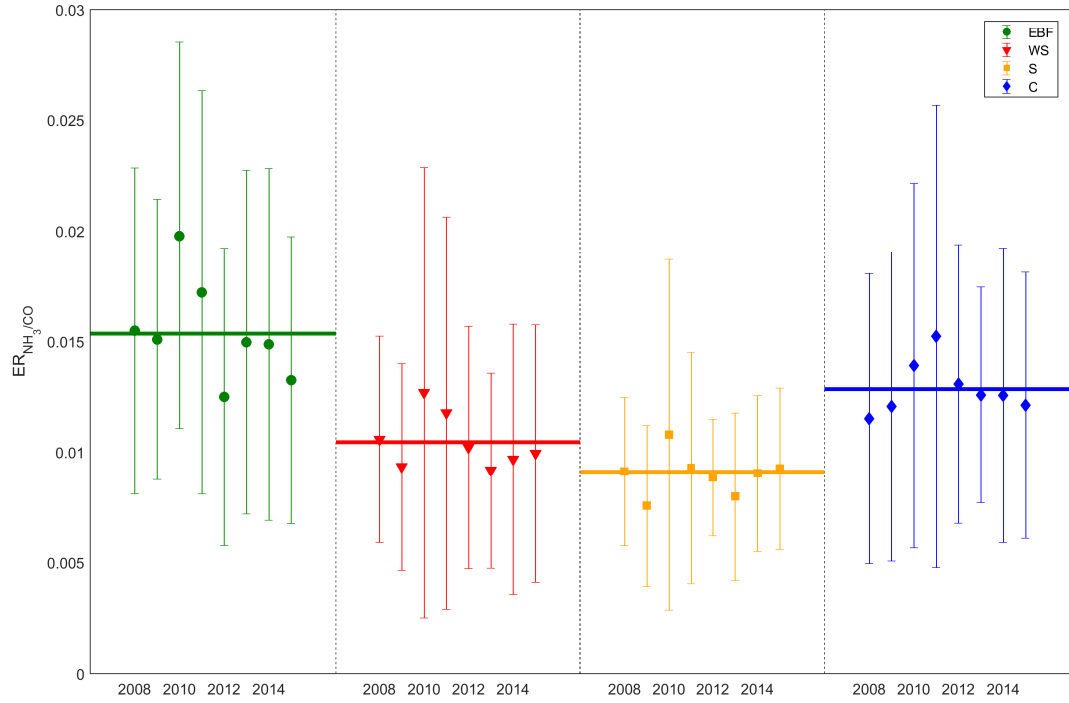


Figure 6. Mean $ER_{NH_3/CO}$ averaged by biome and by year (2008–2015) from the yearly $ER_{NH_3/CO}$ calculated for the pixels selected in section 2.3. The error bar is the 1-sigma standard deviation around the mean. The solid line represents the 8-years average per biome.

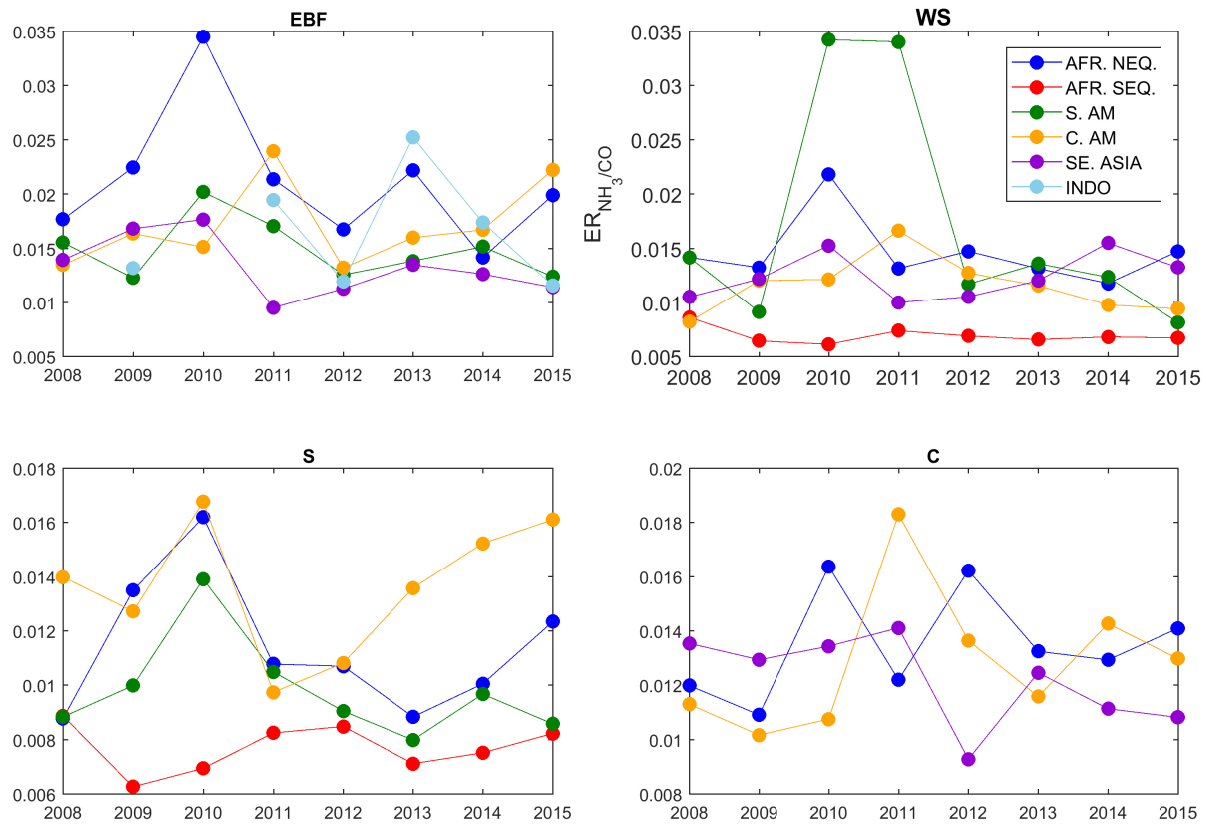


Figure 7. Mean biome-specific $ER_{NH_3/CO}$ averaged by year (2008–2015) and by region (colored dots and lines) from the yearly $ER_{NH_3/CO}$ calculated for the pixels selected in section 2.3. From top-left to bottom-right: EBF, WS, S and C.

Table 1. $ER_{NH_3/CO}$ reported in the literature for different regions and biomes. $ER_{NH_3/CO}$ calculated in this study are given as well.

Source	NC Africa ^a	SC Africa ^b	S America ^c
Luo et al. (2015) - TES	$14-23 \times 10^{-3}$	- ^d	15×10^{-3}
Luo et al. (2015) - GEOS-Chem	$8-17 \times 10^{-3}$	$14-16 \times 10^{-3}$	11×10^{-3}
<i>This study</i>	$11-23 \times 10^{-3}$	$7-8 \times 10^{-3}$	$10-19 \times 10^{-3}$
Source	Savanna	Tropical Forest	Cropland
Andreae and Merlet (2001)	15.2×10^{-3}	20.5×10^{-3}	23.3×10^{-3}
Bertschi et al. (2003)^e	19.9×10^{-3}	-	-
Sinha et al. (2003)	7×10^{-3}	-	-
Yokelson et al. (2003)	$6.5-7 \times 10^{-3}$	-	-
Christian et al. (2007)	12.8×10^{-3}	-	-
Akagi et al. (2011)	13.6×10^{-3}	23.6×10^{-3}	35.0×10^{-3}
Wooster et al. (2011)	$8-35 \times 10^{-3}$	-	-
Yokelson et al. (2011)	9.9×10^{-3}	46.8×10^{-3} ^f	29.1×10^{-3}
Kaiser et al. (2012)	24.3×10^{-3}	15.2×10^{-3}	28.6×10^{-3}
Smith et al. (2014)	13.3×10^{-3}	-	-
<i>This study</i>	$7-19 \times 10^{-3}$	$14-23 \times 10^{-3}$	$11-14 \times 10^{-3}$

^{a,b,c} NC Africa=North-Central Africa, SC Africa=South-Central Africa, S America=South America

^d too low correlation coefficient

^e for smoldering logs

^f Tropical dry forest